



Evaluation of high-resolution meteorological data products using flux tower observations across Brazil

Jamie R. C. Brown¹, Ross Woods¹, Humberto Ribeiro da Rocha², Debora Regina Roberti³, Rafael Rosolem^{1,4}

5 ¹School of Civil Aerospace and Design Engineering, University of Bristol, Bristol, BS8 1TR, UK

²Instituto de Astronomia, Geofísica, e Ciências Atmosféricas, Universidade de São Paulo, São Paulo, 13083-8782, Brazil

³Departamento de Física, Universidade Federal de Santa Maria, Santa Maria, 97105-900, Brazil

⁴ Cabot Institute for the Environment, University of Bristol, Bristol, BS8 1UH, UK

10 *Correspondence to:* Jamie Brown (Jamie.Brown@Hydrology.uk.com), Rafael Rosolem (Rafael.Rosolem@bristol.ac.uk)

Abstract. In the past decade, the scientific community has seen an increase in the number of global hydrometeorological products. This has been possible with efforts to push continental and global land surface modelling to hyper-resolution applications. As the resolution of these datasets increase, so does the need to compare their estimates against local in-situ measurements. This is particularly important for Brazil, whose large continental scale domain results in a wide range of 15 climates and biomes. In this study, high-resolution (0.1 to 0.25 degrees) global and regional meteorological datasets are compared against flux tower observations at 11 sites across Brazil (for periods between 1999-2010), covering Brazil's main land cover types (tropical rainforest, woodland savanna, various croplands, and tropical dry forests). The purpose of the study is to assess the quality of four global reanalysis products [ERA5-Land, GLDAS2.0, GLDAS2.1, and MSWEPv2.2] and one regional gridded dataset developed from local interpolation of meteorological variables across the country [Brazilian National 20 Meteorological Database (referred here as BNMD)]. The surface meteorological variables we considered were precipitation, air temperature, wind speed, atmospheric pressure, downward shortwave and longwave radiation, and specific humidity. Data products were evaluated for their ability to reproduce the daily and monthly meteorological observations at flux towers. A ranking system for data products was developed based on the mean squared error (MSE). To identify the possible causes for these errors, further analysis was undertaken to determine the contributions of correlation, bias, and variation to the MSE. 25 Results show that, for precipitation, MSWEP outperforms the other datasets at daily scales but at a monthly scale BNMD performs best. For all other variables, ERA5-Land achieved the best ranking (smallest) errors at the daily scale and averaged the best rank for all variables at the monthly scale. GLDAS2.0 performed least well at both temporal scales, however the newer version (GLDAS2.1) was an improvement of its older version for almost every variable. BNMD wind speed and GLDAS2.0 shortwave radiation outperformed the other datasets at a monthly scale. The largest contribution to the MSE at the daily scale 30 for all datasets and variables was the correlation contribution whilst at the monthly scale it was the bias contribution. ERA5-Land is recommended when using multiple hydro-meteorological variables to force land-surface models within Brazil.



1 Introduction

In regions that lack high-density meteorological monitoring networks or have sporadic historical observations, gridded weather products provide valuable historical references to aid studies for many purposes, including water resources (Syed et al., 2008; 35 Vissa et al., 2019), flood forecasting and heatwaves (Miralles et al., 2019), prediction of vegetation dynamics and agricultural yields (Tian et al., 2019), and climate change impacts (Wagner et al., 2007; Dullaart et al., 2019; Terzaghi et al., 2020; Xi et al., 2021). These products provide a method to integrate available weather station data both temporally and spatially 40 consistently, whilst taking into consideration factors of influence such as topography, prevailing winds, and distance (Thornton et al., 2021). They are becoming more readily available worldwide and are helping with regional to global applications where ground-based observations are not available or more consistent temporally extensive datasets are needed (e.g., Soti et al., 2010; Hughes and Slaughter, 2015; Gampe 2017). However, limitations in the forcing data can result in disinformation in data which can lead to incorrect conclusions (Beven, 2011; Kauffeldt et al., 2013) and therefore the validation of such products is key for the performance of models. Comparison studies between these products and ground-based observations over the study area is one way to validate and determine its reliability and suitability.

45 New efforts are being made to validate global data products for important hydrological applications. For example, Sikder et al. (2019) tested three GLDAS versions and ERA-Interim/Land products over South and Southeast Asia (the Ganges-Brahmaputra-Meghna and Mekong River basins) against discharge observations to determine which product better describes the system. Gebrechorkos et al. (2020) used rainfall observations to analyse the competency of two gridded high-resolution datasets to detect climate variability and droughts across East Africa, whilst Weber et al. (2021) compared multiple gridded 50 products against an Alpine observation centre to determine their capability for snow hydrological modelling.

The selection of a gridded product is based on its suitability for long-term hydrological applications, which require consistent meteorological forcing data spanning over 20 years or more. However, it is also important to compare products in data-sparse areas whenever ground-based observations are available to provide insight on the local-to-regional uncertainties associated with the product. Higher uncertainties potentially affect the ability to prepare for climate events by local or regional institutions. 55 Furthermore, increased data in areas with scarce availability strengthens model representation of Earth system processes (IPCC, 2012).

With centers such as the European Centre for Medium Range Forecasts (ECMWF), and the National Aeronautics and Space Administration (NASA) Goddard Space Flight Center (GSFC)'s Global Modelling and Assimilation Office (GMAO) producing openly available high-resolution global gridded products using different techniques it can be difficult to know which 60 products may be better suited for each application. Different products may excel in some areas over others due to the nature of their interpolation/reanalysis method and the ground observations used. For example, products like the ECMWF's ERA5-Land (Muñoz-Sabater, 2019) are developed from blending observations with past short-range weather forecasts rerun with modern weather forecasting models to produce many land-surface flux variables. MSWEPV2.2 (Beck et al. 2019), however, focuses only on precipitation data and combines satellite remote sensing data with multiple sources of reanalysis products,



65 then bias corrects and weights between multiple nearby observation gauges. Multiple studies have been undertaken showing that different products provide contrasting results depending on the environment or climate in question (Decker et al., 2012; Wang and Zeng, 2012; Sikder et al., 2019; Beck et al., 2021).

The global distribution of weather stations tends to be biased toward populated areas leaving large areas underrepresented (Viana et al., 2021). For example, in a study in 2018 utilised 11,427 rain gauges covering Brazil, and the Amazon basin, which 70 contains 70% of the country's freshwater, has the lowest density of gauges with only 199 in the entire state of Amazonas (Filho et al., 2018). Places of ecological importance such as large forests and savannahs like the Amazon and Cerrado influence the hydrological cycle through a variety of factors including biodiversity, vegetation dynamics, and root distribution (Oliveira et al., 2005; Diaz et al., 2007, Bonal et al., 2016, Coe et al., 2016) despite low meteorological station density. To combat this, meteorological observation data from measurement stations such as flux towers and eddy covariance stations provide valuable 75 insight into representative areas across the country.

Fluxnet is an international network of flux towers where eddy covariance techniques are used to measure energy, water and carbon fluxes between the biosphere and atmosphere (Balocchi et al. 2001), however, their distribution is highly biased towards North America and Europe. This network of towers has provided opportunities to validate gridded products over data rich areas such as North America (Decker et al., 2012) and China (Wang and Zeng, 2012) but less has been done in 80 comparatively data poor areas such as South America. The tower sites in this study are part of Fluxnet but no attempt thus far has been made to comprehensively evaluate gridded products using these locations.

Here we compare the accuracy of one regional and four global high-resolution gridded meteorological products over 11 ecologically different flux tower sites spanning multiple biomes across Brazil (Fig. 1, Table 1). Specifically, the study aims to answer four questions. Firstly, which high-resolution gridded product is the most accurate overall when compared with local 85 observation data? Secondly, which product is the most accurate when compared with observation data each variable considered? Thirdly, what are the dominant types of error associated with each product when compared to observation data? And finally, how do these errors vary by location and time of year?

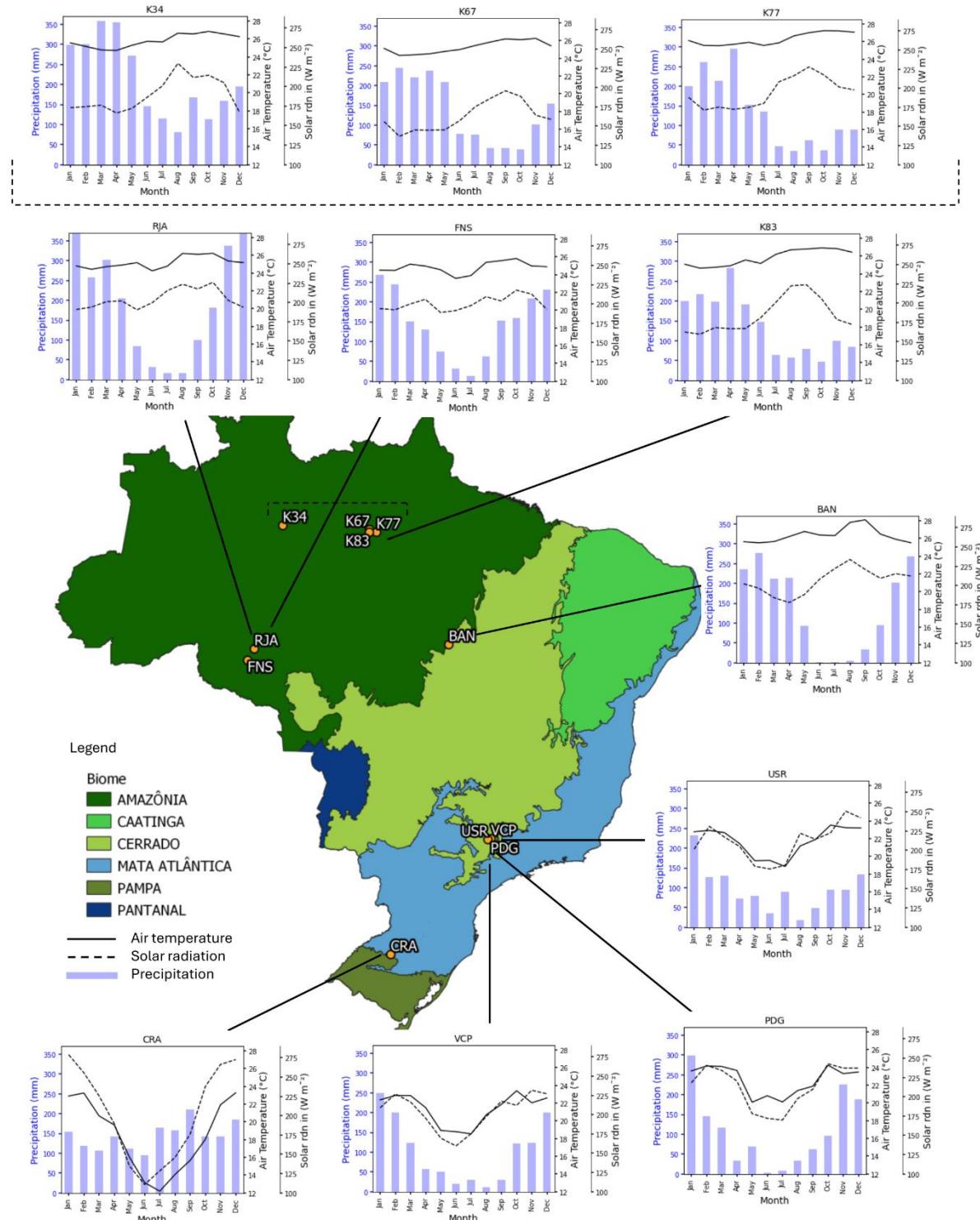


Figure 1: Map of Brazil spilt into its six major biomes showing the location of the study sites analysed in this paper and their measured average monthly precipitation, air temperature, and incoming irradiance.



Table 1 Summary of the abbreviated names, lat, lon, location, period of observation, vegetation type, elevation (abs), and references for the 11 flux tower sites analysed in this paper. All sites were provided directly by the principal investigator. *

Name	Lat (°N)	Lon (°E)	Location (State)	Start	End	Land Cover	Elevation (m)	Reference
K34	-2.6091	-60.2093	Manaus (AM)	2000	2006	Tropical Rainforest	130	Araujo et al. 2002
K67	-2.8567	-54.9589	Tapajos (PA)	2002	2006	Tropical Rainforest	130	Rice et al. 2004
K83	-3.0117	-54.9707	Tapajos (PA)	2000	2004	Tropical Rainforest	130	Goulden et al. 2004
K77	-3.0119	-54.5365	Tapajos (PA)	2000	2005	Cropland - Pasture	130	Sakai et al. 2004
RJA	-10.083	-61.9309	Jaru (RO)	2000	2002	Tropical Dry Forest	191	Andreae et al. 2002
FNS	-10.774	-62.3374	Ji-Parana (RO)	2000	2004	Cropland - Pasture	306	Kirkman et al. 2002
BAN	-9.8244	-50.1591	Araguaina (TO)	2003	2017	Woodland Savanna	120	Borma et al. 2009
PDG	-21.621	-47.63	Luis Antonio (SP)	2000	2017	Savanna	690	Rocha et al. 2002
USR	-21.637	-47.7903	Luis Antonio (SP)	2005	2008	Cropland - Sugarcane	552	Cabral et al. 2013
VCP	-21.583	-47.602	Ribeirao Preto (SP)	2006	2009	Cropland (Eucaliptus)	761	Cabral et al. 2010
CRA	-28.603	-53.6736	Cruz Alta (RS)	2009	2014	Cropland (soybean)	432	Webler et al. 2012

*The site names are abbreviations used throughout this paper. K stands for Kilometer followed by a number which the name of the access road to the site within the Amazon. The other sites are abbreviations also regarding their location listed here: CRA – Cruz Alta, FNS – Fazenda Nossa Senhora, RJA – Rebio Jaru, PDG – Pé de Gigante, VCP – Votorantim, Fazenda Cara Preta, BAN – Banana Island, USR – Usina Santa Rita



2 Datasets

2.1 In-situ observations

Meteorological data were obtained from the 11 sites across Brazil and are presented in Fig. 1 and Table 1. Data collected from 95 these sites range from four to seven years, and do not always cover the same period. This study focuses on the main hydrometeorological variables utilised in forcing land-surface models, creating surface flux estimates and are used in calculating evapotranspiration used to drive many hydrological models. Variables were included so that reference evapotranspiration could be calculated according to the standard FAO methodology (Allen et al., 1998). For each site these 100 were selected based on the quality and availability and, where needed, converted for uniformity and comparability (see section 3, Methodology). Table 2 provides the recorded observation variables available at each site, which could be converted to the desired variables to be analysed using standard FAO methods. The desired variables to be analysed were air temperature, precipitation, wind speed, air pressure, longwave and shortwave radiation, and specific humidity (Table 2 and Table 3).

Table 2 Availability of each variable for each flux tower site. The meteorological variables are; *ws*, wind speed, *ta*, air 105 temperature, *press*, atmospheric pressure, *rgs*, short wave radiation, *par*, photosynthetically active radiation (used to calculate *rgs*), *rgl*, long wave radiation, *prec*, precipitation, *ee*, vapour pressure, *dpt*, dew point temperature, and *RH*, relative humidity. *ee*, *dpt* and *RH* variables were used to calculate specific humidity.

Site	Meteorological variables									
	<i>ws</i>	<i>ta</i>	<i>press</i>	<i>rgs</i>	<i>par</i>	<i>rgl</i>	<i>prec</i>	<i>ee</i>	<i>dpt</i>	<i>RH</i>
K34	X	X	X	X		X	X	X		
K67	X	X	X		X		X	X		
K77	X	X	X	X		X	X			X
K83	X	X	X	X		X	X	X		
CRA	X	X	X	X			X			X
FNS	X	X	X	X		X	X	X		
RJA	X	X	X	X		X	X	X		
PDG	X	X	X	X			X	X		
VCP	X	X	X	X			X			X
BAN	X	X	X	X			X	X		
USR	X	X	X	X			X			X



2.2 Regional Products

Efforts have been made to produce high-resolution datasets through interpolation of weather stations (Xavier et al., 2016). The meteorological station network across Brazil varies spatially and temporally with few data available before 1980. There has been a steady increase in weather stations and rain gauges over the last 40 years but with heavy bias towards stations closer to 115 densely populated areas such as São Paulo and Rio de Janeiro (Alvares et al., 2013; Filho et al., 2018). These biases bring the quality of meteorological datasets under scrutiny and a strong need for validation especially over the more data poor areas. Gridded data products were selected based on their open access availability, a spatial resolution of 0.25 x 0.25 degrees or finer, and a daily or sub-daily temporal resolution (Table 3) referred to here on as high-resolution.

2.2.1 Brazilian National Meteorological Gridded Database (BNMD)

120 A high-resolution gridded dataset developed from local interpolation of meteorological variables across Brazil was released in 2016 (the Brazilian National Meteorological Database, referred here as BNMD) (Xavier et al. 2016). The data were collected from 3625 rain gauges and 735 weather stations over the period of 1980-2013 and quality control procedures were performed to identify outliers based on Liebmann and Allured (2005). A lack of previous reviews due to the novelty of this dataset combined with the rapid increase in stations/gauges over the 30-year period of data that has been made available, raises 125 questions about its reliability, particularly over less data rich areas such as the Amazon.

2.2.2 Global Land Data Assimilation System (GLDAS) 2.0 and 2.1

In 2004, NASA-GSFC and NCEP released a reanalysis data product called Global Land Data Assimilation System (GLDAS) (Rodell et al. 2004). Since then, GLDAS has been reprocessed leading to the updated release of GLDAS2.0 in November 2019 and GLDAS2.1 in January 2020. GLDAS2.0 data are products of the new NOAH-3.6 LSM forced using the Princeton 130 meteorological forcing dataset (Sheffield et al. 2006) producing a dataset from 1948 – 2014. GLDAS2.1 is a direct update from GLDAS-1 where NOAH-3.6 LSM is forced with combined forcing data including Global Precipitation Climatology Project (GPCP) version 1.3 produced by NOAA with available data from 2000-present.

2.2.3 ECMWF Reanalysis 5-Land (ERA5-Land)

In 2019, the European Centre for Medium Weather Forecasting (ECMWF) released ERA5-Land (an upgraded form of ERA-135 Interim) providing a higher resolution global land-based dataset from 1981-present (0.1 x 0.1 degrees, hourly) (Muñoz-Sabater, 2019) generated using Copernicus Climate Change Service Information. The production of ERA5-Land is the result of the tiled ECMWF Scheme for Surface Exchanges over Land incorporating land surface hydrology (H-TESSEL). The recent release sees it benefit from over a decade of developments in 4D-VAR data assimilation, core dynamics, and model physics relative to GLDAS and ERA-Interim. As it integrates a wide array of global observation data sources, employs advanced data



140 assimilation techniques, and benefits from continuous improvements the quality would be expected to be higher than that of new regional datasets such as ones produced by Xavier et al. (2016).

2.2.4 Multi-Source Weighted-Ensemble Precipitation v2.2 (MSWEPv2.2)

Another recent dataset that garnered significant attention is the Multi-Source Weighted-Ensemble Precipitation, version 2.2 (MSWEP V2.2) (Beck, et al., 2019). Although only precipitation data, it provides the highest temporal (3-hourly) and spatial 145 (0.1 degrees) resolution based on gauges, satellites, and reanalysis with distributional bias corrections. The dataset merges multiple observation, satellite and reanalysis data across the globe and its predecessors have proven to provide reliable estimates for precipitation patterns globally (Baez-Villanueva et al. 2018; Moreira et al. 2018; Alijanian et al. 2019; Xu et al. 2019).

150 Table 3 provides a summary of the gridded products used in this study with information about time periods covered, temporal and spatial resolution, the meteorological variables accessed and their references.

Table 3 Summary of gridded data products used in this study*

* Refer to Table 2 for abbreviations. q refers to specific humidity. Note: temporal resolution was converted to the most coarse of the datasets – Daily. Monthly 155 values were also generated.

Data Descriptor	Data Source	Variables accessed	Periods	Temp. res.	Spatial res.	Reference
BNMD	BNMD	Prec, maxTa, minTa, rg, press, ws, RH	1980-2017	Daily	0.25x0.25	Xavier et al. 2016
ERA5-Land	ERA5-Land	Prec, Ta, rg, rgl, press, ws, dpt	1981-2019	Hourly	0.1x0.1	Muñoz Sabater, 2019
GLDAS2.0	GLDAS_NOAH25_3H 2.0	Prec, Ta, rg, rgl, press, ws, q	1948-2014	3-hourly	0.25x0.25	Rodell et al. 2004
GLDAS2.1	GLDAS_NOAH25_3H 2.1	Prec, Ta, rg, rgl, press, ws, q	2000-2019	3-hourly	0.25x0.25	Rodell et al. 2004
MSWEPv2.2	MSWEP_v2.2_sh	Prec	1979-2017	3-hourly	0.1x0.1	Beck et al. 2019

3. Methodology

This section describes the data manipulation necessary that enabled us quantification of the differences between the gridded products and observations.

3.1 Quality control

160 Flux tower data frequently contain gaps or periods with suboptimal data quality due to a variety of reasons (e.g. sensor malfunction, drifting, calibration errors, power supply issues). To address this, the selection of variables for each site was guided by their data availability, consistency, and a requirement for completeness, with more than 80% data coverage achieved after infilling (Jung et al., 2024). Furthermore they were analysed for trends to identify potential measurement drifts caused



165 by instrumentation. Initial quality control and gap-filling procedures had already been conducted by the principle site investigator. Despite this, remaining errors related to faulty instrumentation were removed from the dataset. Gaps in the seven meteorological variables were subsequently filled using linear regression, prioritising variables from the same site that exhibited strong correlations ($R^2 > 0.8$), with the most robust correlations being utilised first.

3.2 Temporal averaging

170 Flux tower sites have different recording methods and temporal resolutions. All observation and gridded datasets were converted to the coarsest common temporal resolution, the daily scale, for analysis (BNMD, Table 3). As the gridded datasets have no gaps this was a straightforward forward or backward averaging depending on the variable and averaging method. Two-sample Kolmogorov-Smirnov (K-S) tests were carried out on the observation data for each variable, where full days (24-hours) were used to create daily data and set as the reference distribution. Samples were then tested against this distribution using one less hour each iteration to determine whether samples significantly deviated from the reference sample. 12 hours of 175 data (50% available) in a day was set as the minimum for daily conversion.

A similar averaging method was adopted to convert daily data to monthly. However, due to a shortage of data availability, instead of using 100% of days available in a month as the reference sample, 80% or above was used to acquire a more representative sample to test against. The minimum conversion allowed was 50% of days in a month.

180 Precipitation was summed when converting to daily and monthly. Rainfall does not follow a regular pattern or known distribution, meaning taking anything less than all 24 hours of available data would result in an under-prediction. Therefore, only days with all hourly data available were converted to daily. The same approach was taken converting daily to monthly but, in some cases, resulted in a high loss of data. To conserve data, each site was assessed uniquely looking at the two-sample K-S test results and changes in the mean after using fewer days in the month to convert (i.e., rejected if changes in the means and standard deviations were $>2\%$). A scaling factor was then applied to the monthly total depending on how the percentage 185 of days missing to bring the total up to 100%.

3.3 Wind Speed vertical interpolation

190 The height at which the measurement instruments are located differ at each site. To compare data products to the observation data they are vertically interpolated to the height of the instrument at each site. The BNMD wind speed variable was calculated by interpolating laterally from the nearest Brazilian weather station which records wind speed over grass. The ERA5-Land 10m wind speed product is produced for comparison against SYNOP stations, also above grass. GLDAS 10m wind speed is adjusted down from the model's lowest level to 10m but it is unclear whether this is over grass or different vegetation types. For consistency, the same vertical interpolation method was used for all data products. The wind speeds were interpolated up the log-wind profile using grass as the vegetation type at the height of the WMO weather station standard (30cm) from either 2m or 10m depending on the data product.



195 3.4 Atmospheric Pressure

Atmospheric pressure is not provided as a variable in the BNMD dataset; thus, it was estimated as a single continuous value following the method outlined in the FAO Irrigation and Drainage Paper No. 56 (based on data homogeneity and availability). Incorporating this estimated value allowed for a critical comparison of its performance relative to other measured variables, providing insight into whether the observed variables performed better or worse than a single mean estimate. The atmospheric 200 pressure variable was available for comparison in the ERA5-Land and both GLDAS products.

3.5 Specific humidity

Specific humidity was available for both GLDAS datasets but needed to be calculated for ERA5-Land and BNMD. ERA5-Land water vapour pressure was calculated from the dew point temperature variable and then converted to specific humidity using pressure (Shuttleworth, 2012). For BNMD, the vapour pressure at maximum and minimum temperatures was calculated 205 using the FAO method (Eq. 11, Allen et al. 1998). These were then used with the relative humidity, and estimated pressure to calculate specific humidity using ideal gas laws (Bolton, 1980).

3.6 Decomposition of the mean square error

To quantify the differences the Mean Square Error (MSE) was calculated for each variable and data product against the observation data at each site. The MSE is a single metric with limited explanation about the source of the error, but it can be 210 decomposed into parts to acquire a better understanding of contributions to the error (Gupta et al. 2009) (Eq. 1)).

$$MSE = 2 \cdot \sigma_s \cdot \sigma_o \cdot (1 - r) + (\sigma_s - \sigma_o)^2 + (\mu_s - \mu_o)^2 \quad (1)$$

In Eq. 1, σ_s and σ_o are the standard deviations of the sample (gridded product) and observations, r is the linear correlation 215 between the sample and the observations, and μ_s and μ_o are the means of the sample and observations respectively. Written like this the equation is seen to have three parts. The first term is the correlation contribution to the MSE, the second, the variation contribution or differences in standard deviation, and the third term represents the bias contribution or differences in means. For a clearer visualisation of the results, the individual error source to MSE was scaled to the RMSE magnitude to 220 conserve the units for each variable (Iwema et al., 2017). Whilst the MSE is beneficial when comparing products, quantifying the relative contributions to the MSE provides valuable insights into the reasons behind discrepancies between the observations and the gridded products.

3.7 Performance ranking

The large volume of data, spanning five gridded products, decades of time series, and seven variables across multiple observation sites, presents significant challenges in summarising the results into a coherent structure. For this reason, a ranking

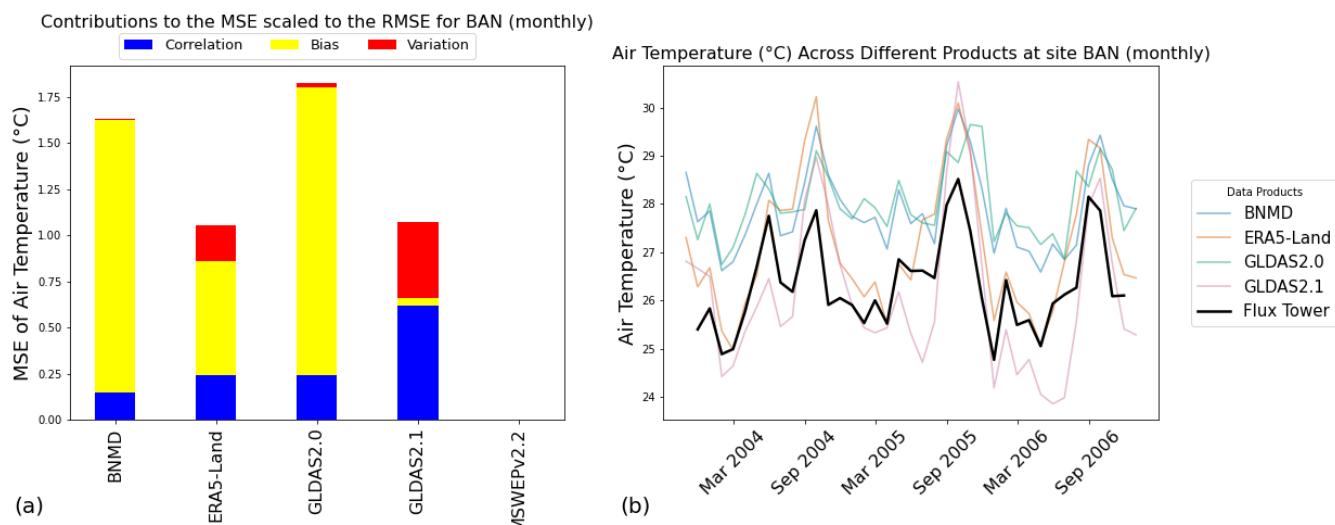


225 system was used to ascertain which data product for each variable performed best due to its simplicity (Brunke et al. 2003). The MSE for each variable was given a rank dependant on how many data products have that variable recorded (for incoming longwave radiation this was 1-3; precipitation, 1-5; and all other variables 1-4, 1 being the best performing/lowest MSE). This was carried out for each site and then the ranks were averaged across all 11 sites to provide a single rank for each variable and each data product. A product with the lowest MSE for a variable over all 11 sites would score a rank of 1. An overall average 230 rank was then also given to each data product which only included ranks of variables that were present for all products. This method of ranking was performed for both the daily and monthly data.

4 Results

In this section, precipitation and air temperature are analysed separately due to their fundamental importance in hydrological and climatological studies. These two variables exert significant influence on ecosystem dynamics and are the more widely 235 used in model validation and environmental monitoring. A more detailed examination is thus warranted. The remaining variables, while important, are discussed collectively for brevity, as they primarily serve to complement the analysis of temperature and precipitation.

We applied our methods to five products at 11 observation sites for seven meteorological variables, at both daily and monthly 240 timescales. An example of the results is shown in Fig. 2, which illustrates monthly air temperature at the BAN site. The visual representation of the MSE components allows for a clearer interpretation of the performance of each dataset. In Fig. 2a, the partial contributions to the MSE are colour-coded and plotted as stacked columns, where the total column height reflects the total MSE, scaled to the RMSE. The column height differences across products facilitate direct comparison of their performance.



245 **Figure 2** Air temperature at site BAN, (a) partial contributions to the MSE for each gridded dataset, and (b) the monthly time series of all gridded datasets for air temperature over the operation period of the flux tower (observation data in bold).



To further understand these results, Fig. 2b offers additional context, illustrating the key sources of error presented in Fig 2a. For example, the large bias contributions for the BNMD and GLDAS2.0 datasets are attributed to consistent overpredictions of observed data with higher mean temperatures. Despite this, the seasonality of both datasets aligns closely with observations, resulting in a lower contribution from variability errors. In contrast, the GLDAS2.1 dataset displays higher variability error, 250 likely due to overprediction during the hot months and underprediction during the cooler periods, though it has a smaller bias contribution due to a similar overall mean with the observation data. The ranking scores in Table 4 complement these visualisations (Fig. 3 and Fig. 4) by providing a comprehensive comparison of the MSE for each dataset and variable, averaged over the 11 flux tower sites and presented alongside an overall ranking across all variables.

4.1 Precipitation

255 Precipitation data from all five gridded datasets were statistically analysed at both daily and monthly temporal resolutions, allowing for a comprehensive comparison across scales. MSWEPv2.2 consistently performed the best at the daily scale, while GLDAS2.0 exhibited the poorest performance at both daily and monthly resolutions. However, when considering monthly precipitation, BNMD demonstrated the best overall performance among the compared products.

Figure 3 and 4 illustrate the partial contributions to the MSE for each dataset at daily (Fig. 3) and monthly (Fig. 4) scales, 260 providing insight into the sources of error across sites. At the daily scale (Fig. 3), the primary contributor to the MSE is the correlation error, particularly at sites such as RJA and K34, where variability errors are also prominent. Interestingly, bias contributes very little to the total MSE on a daily basis. However, this pattern shifts at the monthly scale (Fig. 4), where bias plays a much larger role. For instance, in the northern Amazonian sites (K67, K77, and K83), the ERA5-Land dataset consistently overpredicts monthly rainfall, resulting in large spikes in the bias component of the MSE.

265 Another key observation is the variability contribution, which is notably high at sites RJA and USR. While most datasets capture the total rainfall during dry months, they tend to overpredict rainfall during the wet season, increasing the variability error. Although MSWEPv2.2 generally outperforms the other datasets, it did not perform consistently across all sites. For example, at the daily scale, MSWEPv2.2 had the poorest performance at K83, where timing discrepancies during the wet season led to high correlation errors. Similarly, at the monthly scale, MSWEPv2.2 performed worst at BAN, where 270 overpredictions during peak wet months resulted in higher bias and variability contributions to the MSE.

4.2 Air Temperature

Air temperature was analysed across four of the gridded datasets. ERA5-Land performed best whilst GLDAS2.0 performed least well at both daily and monthly scales (Table 4). The ranking of 1 and 1.09 indicates that ERA5-Land had the lowest MSE when compared with every other dataset across all sites at the monthly scale and all sites except one (CRA) at the daily scale 275 respectively. The monthly BNMD dataset performed equally as poorly as GLDAS2.0 meaning the ranking system is unable to identify which dataset reflects the in-situ observations least well. Correlation error has the largest contribution to the MSE for the daily datasets with the bias contribution also having some influence (Fig. 3). However, as found with precipitation, at



the monthly scale, the greatest source of error shifts across all datasets and sites from the correlation contribution to the bias contribution (Fig. 4).

280 Both the BNMD and GLDAS2.0 datasets consistently overpredict air temperature explaining the bias contributions for both monthly and daily datasets. Although performing well overall, the monthly GLDAS2.1 dataset had the largest variability contributions which are explained by overpredicting temperatures in the hotter months and underpredicting them in the cooler months (K34, K67, K77 & BAN). The ERA5-Land dataset followed the mean of the observation data most closely but varied in either overpredicting or underpredicting temperature at different sites.

285 **Table 4 Overall ranks for MSE.** MSE is taken per variable per site and ranked (Table S1-S12). Ranks are then averaged for all sites to produce an overall rank for daily and monthly data. Both the lowest (bold) value (i.e. best performance) and highest (italics) value (i.e., worst performance) in each row are identified. Dashed cells (-) indicate no data available. Ranking for precipitation incorporates the fifth dataset MSWEPv2.2 in its calculation.

	Daily				
	BNMD	ERA5-Land	GLDAS2.0	GLDAS2.1	MSWEPv2.2
Wind Speed (m s ⁻¹)	2.55	1.64	3.27	2.55	-
Air Temperature (°C)	3.36	1.09	<i>3.64</i>	1.91	-
Pressure (hPa)	2.18	1.82	2.82	<i>3.18</i>	-
Shortwave rdn in (W m ⁻²)	3.55	1.36	2.36	2.73	-
Longwave rdn in (W m ⁻²)	-	1	2.2	2.8	-
Precipitation (mm)	2.04	2.62	<i>3.42</i>	2.33	1.6
Specific Humidity (kg kg ⁻¹)	<i>4</i>	1.18	2.45	2.36	-
Average Rank (exc. Longwave)	2.95	1.62	2.99	2.51	-

	Monthly				
	BNMD	ERA5-Land	GLDAS2.0	GLDAS2.1	MSWEPv2.2
Wind Speed (m s ⁻¹)	1.91	2	3.36	2.73	-
Air Temperature (°C)	3.27	1	3.27	2.45	-
Pressure (hPa)	2.18	1.82	2.82	<i>3.18</i>	-
Shortwave rdn in (W m ⁻²)	2.73	<i>3</i>	1.91	2.36	-
Longwave rdn in (W m ⁻²)	-	1	2.6	2.4	-
Precipitation (mm)	1.6	3.13	3.2	2.26	1.82
Specific Humidity (kg kg ⁻¹)	<i>4</i>	1.27	2.36	2.36	-
Average Rank (exc. Longwave)	2.62	2.04	2.82	2.56	-

4.3 Other Meteorological Variables

290 Besides precipitation and temperature, five other meteorological variables were analysed: wind speed, pressure, downward shortwave and longwave radiation fluxes, and specific humidity. Among these, wind speed exhibited the poorest performance in GLDAS2.0 across both temporal resolutions, whereas ERA5-Land demonstrated the highest accuracy at the daily scale and BNMD at the monthly scale. Substantial bias errors are evident at site FNS across all datasets, with gridded datasets underpredicting observation data by means ranging 45–75%. At site CRA, a high degree of variability and bias contributes to



295 the MSE, with datasets underestimating observed values from 2009 to 2013. However, from 2013 to 2014, the observed wind speed declines uncharacteristically, leading to an overestimation by the gridded products.

ERA5-Land proved to have the lowest MSE on average whilst GLDAS2.1 performed least well at both temporal resolutions when analysing pressure. It is worth noting that the ranking did not change between daily and monthly scales for pressure as performance consistency was unaffected between daily and monthly datasets. The errors associated with pressure are heavily

300 dominated by the bias contribution (Fig. 3 and Fig. 4). Contributions to the variability error are visible for BNMD as pressure was estimated using the elevation of the site using the standard FAO method (Allen et al. 1998) and kept as a constant figure. BNMD's relatively low MSE when compared to other datasets tells us that estimating a single value for pressure can sometimes more accurately reflect the observation data. The large consistent biases at sites K67 and RJA are due to an overprediction and underprediction of 23 hPa and 19 hPa on average, respectively, for all datasets.

305 ERA5-Land performed best again at the daily scale for the variable, downward shortwave radiation whilst BNMD performed least well. Surprisingly, ERA5-Land performed least well at the monthly scale while GLDAS2.0 performed best. The correlation contribution to the MSE dominated across all sites at both daily and monthly scales, but as temporal resolution decreases, so does the correlation contribution resulting in lower overall MSEs. Large bias contributions are evident at site K67 over both temporal resolutions as the gridded datasets consistently overpredict the observation data by around 40 W m^{-2} .

310 The observation data tends to have a downwards trend over the entire recording period resulting in an increased bias towards the end of the time series when comparing to the gridded products.

Only three of the gridded data products and five sites recorded measurements of downward longwave radiation leading to its exclusion in the overall ranking across all variables in Table 4. ERA5-Land performed best at all sites across both time scales whilst GLDAS2.1 and 2.0 performed least well at the daily and monthly scales, respectively. All datasets tend to underpredict

315 at every site, with contributions to all three components of the MSE visible at both time scales. However, the scale of the errors is not large, ranging between 1-7% error across the spread of the data.

With regards to specific humidity, ERA5-Land outperformed the other gridded datasets again whilst BNMD had the weakest performance at both time scales. BNMD's large biases are due to the estimation of vapour pressure from the minimum and maximum temperatures and a constant estimate for pressure (see Section 3.4). Biases associated with air temperature for

320 BNMD can therefore be expected to be seen in specific humidity. Similarly, the variability contributions to BNMD MSEs found at sites PDG, CRA, VCP and USR are associated with the variability errors in pressure as this was also utilised in the calculation.

The ranks were averaged for all shared variables for BNMD, ERA5-Land, GLDAS2.0 and GLDAS2.1 across both time scales. ERA5-Land performed best on average whilst GLDAS2.0 performed least well at both monthly and daily scales.

325 4.4 Seasonality in Errors

Errors throughout the year can change if the datasets fail to capture the correct range of seasonality. For example, dry seasons may have low errors in precipitation because the mean rainfall will be closer to 0. Figure 5 shows this behaviour across almost



all sites for the best performing precipitation dataset, MSWEPv2.2. Similarly, biases may occur if datasets overpredict temperatures in the warmer seasons as seen in Fig. 6 at sites, BAN and FNS. It is clear from Fig. 1 that seasonality changes 330 with latitude and that sites located further south have a higher range of temperature between seasons. This increased seasonality helps explains the relatively large errors seen at sites CRA and USR in Fig. 6. Comparing the errors spread over of the year between datasets helps us determine which ones best predict the seasonality. For example, take BNMD air temperature, the correlation component's contribution to the MSE increases in the summer months the more southerly the site, suggesting there 335 is a weakness in the datasets ability to predict seasons. Further graphical representations of this can be found in the supplementary material.

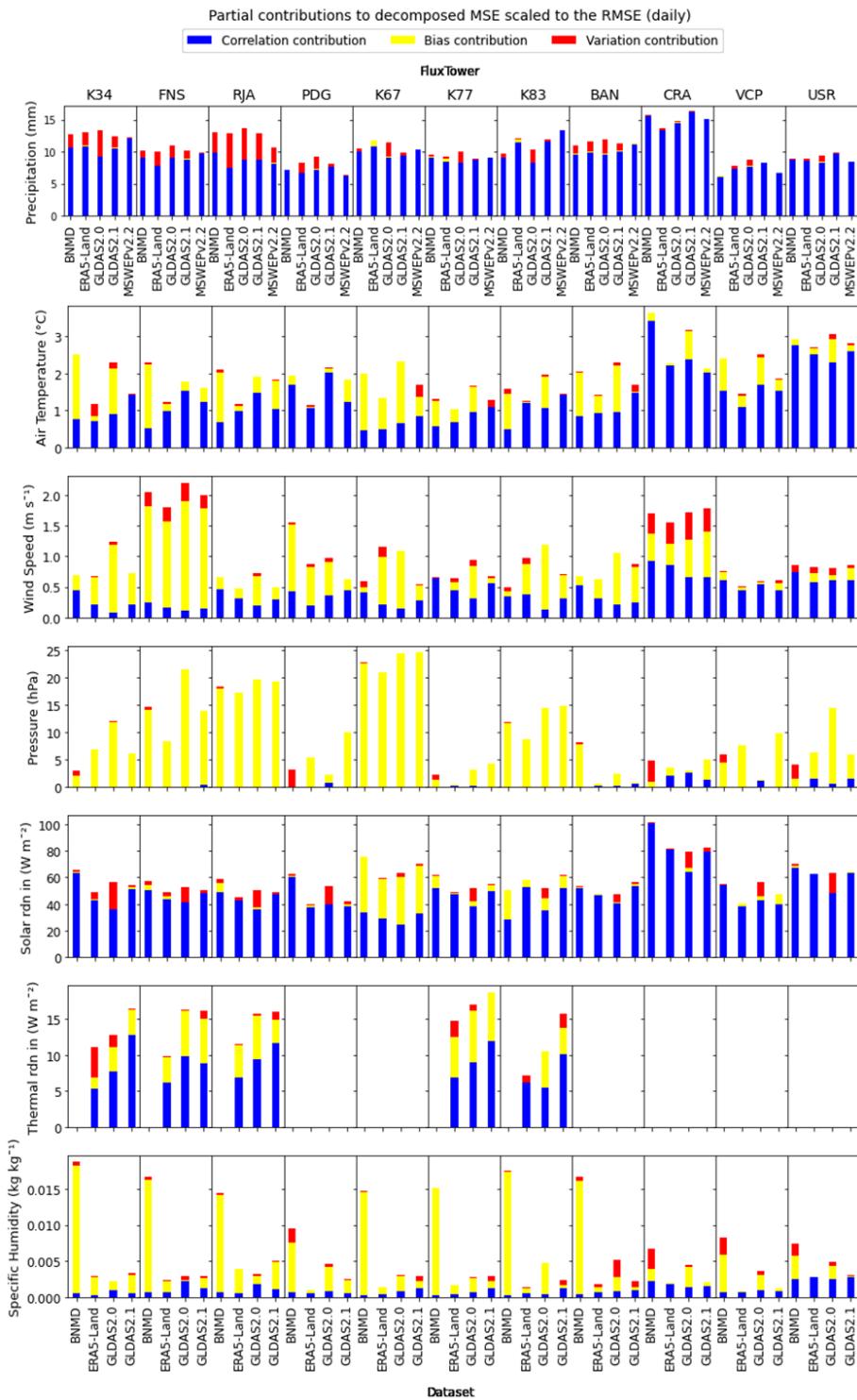


Figure 3 Partial contributions to the MSE between each observation site and gridded dataset (x-axis) for each variable (y-axis) at the daily scale. Precipitation includes MSWEPv2.2 errors.

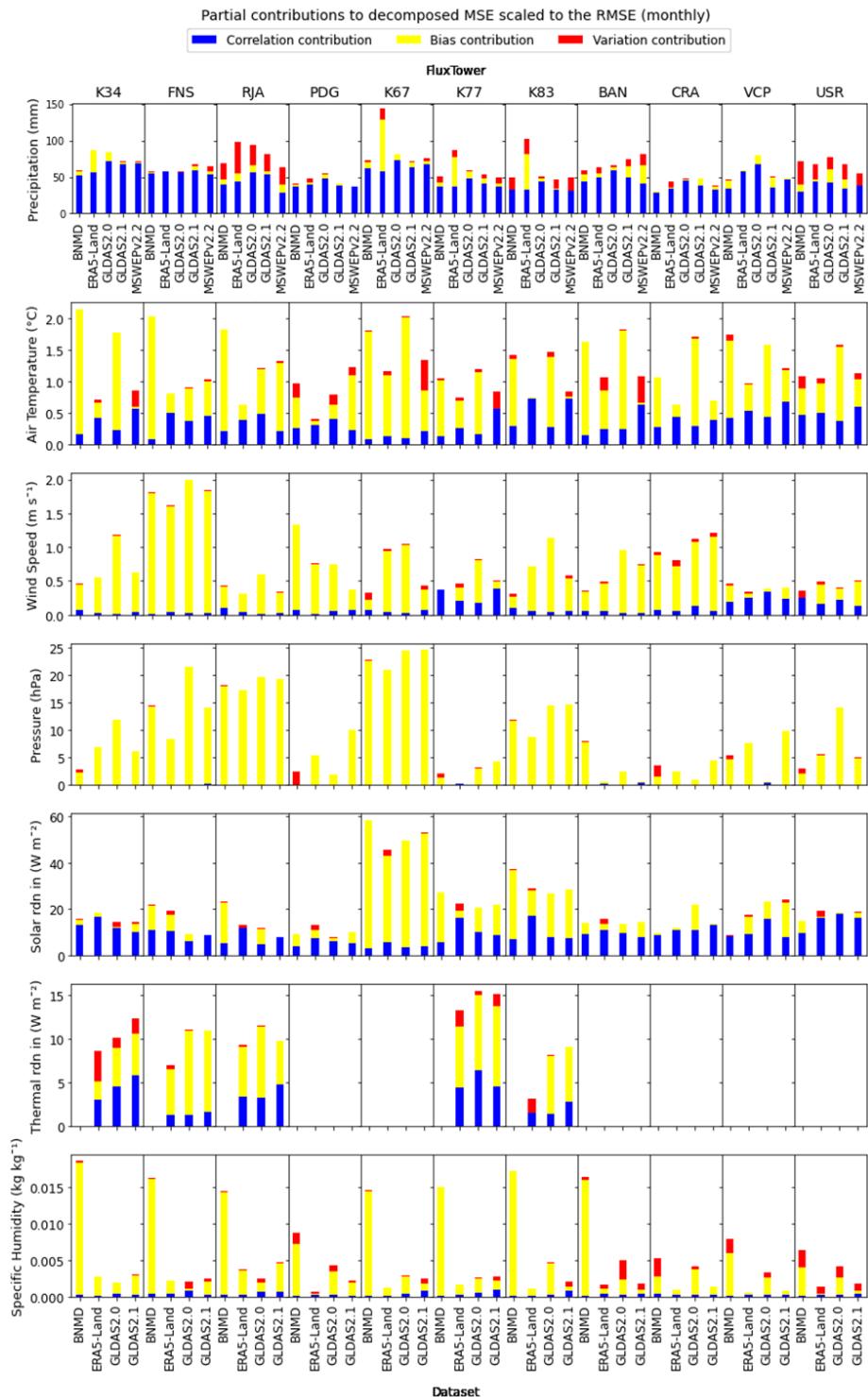


Figure 4 Partial contributions to the MSE between each observation site and gridded dataset (x-axis) for each variable (y-axis) at the monthly scale. Precipitation includes MSWEPv2.2 errors.



All MSWEPv2-2 Precipitation (mm) Partial contributions to decomposed MSE scaled to the RMSE (daily data)

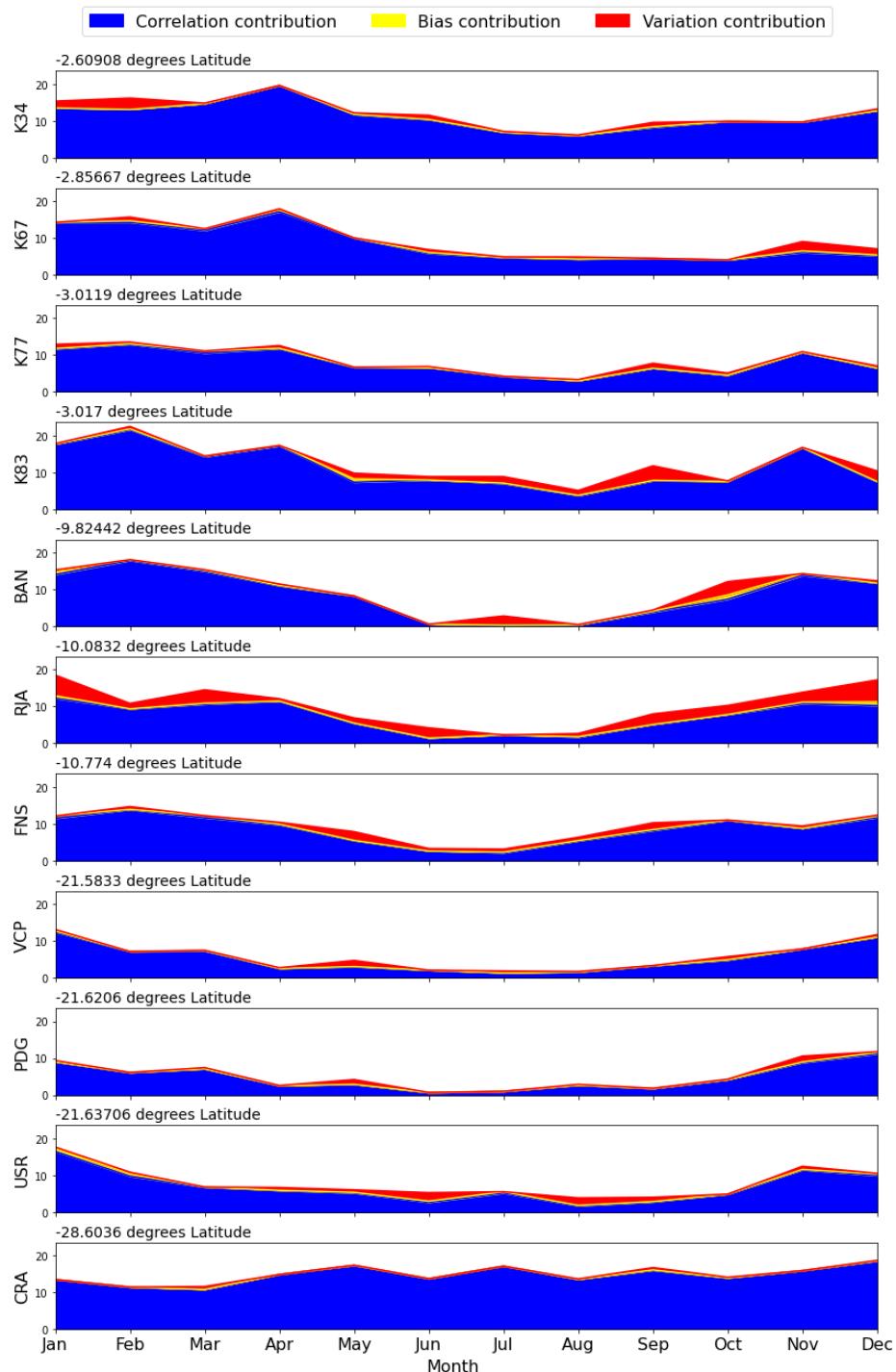


Figure 5 Partial contributions to the MSE averaged by month over all operational observation years for MSWEPv2.2 precipitation across all sites. Sites are in descending order from distance from equator.

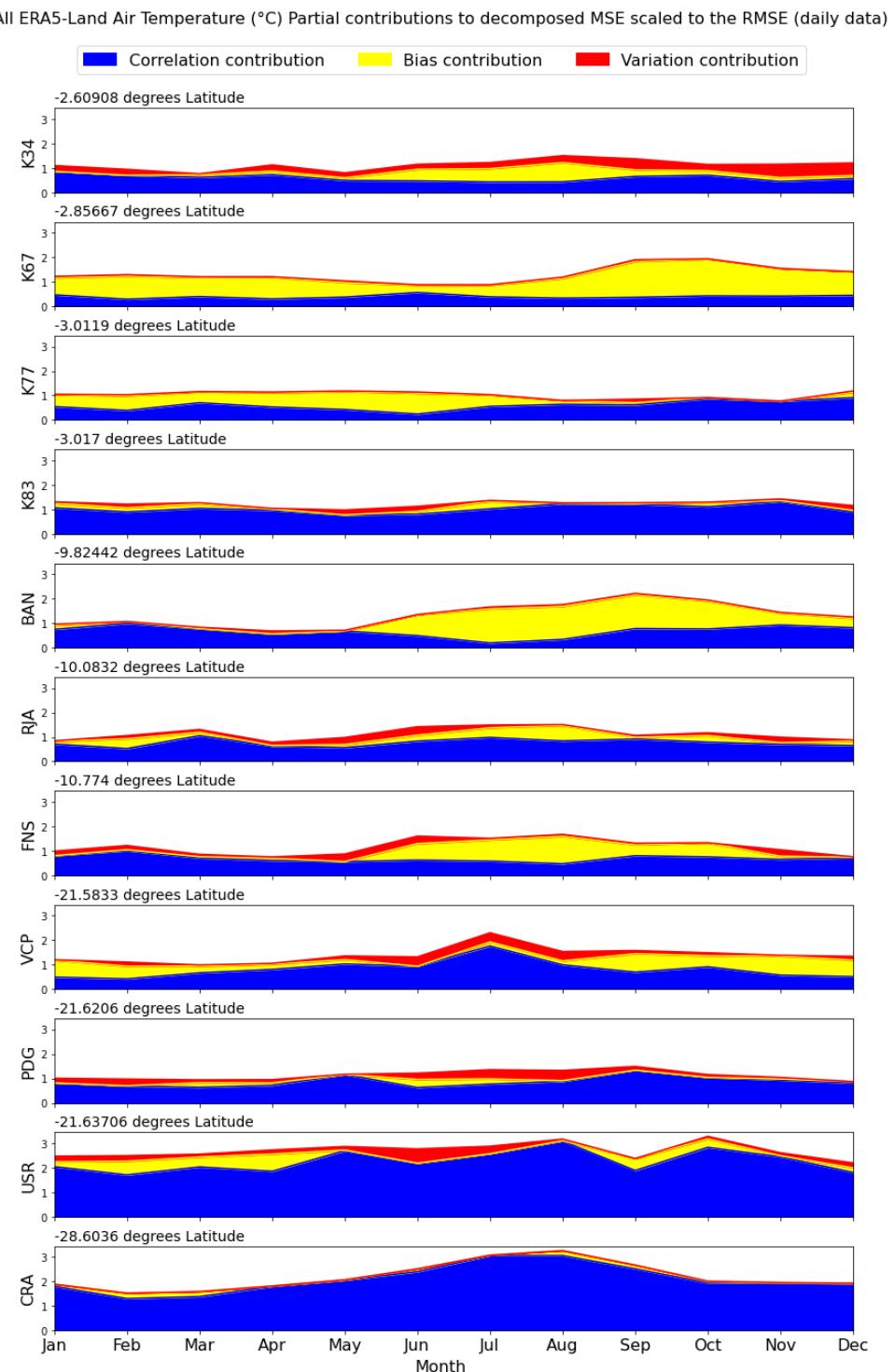


Figure 6 Partial contributions to the M_{SE} averaged by month over all operational observation years for ERA5-Land air temperature across all sites. Sites are in descending order from distance from equator.



345 **5. Discussion**

Five gridded data products (BNMD, GLDAS2.0, GLDAS2.1, ERA5-Land, and MSWEPv2.2) were evaluated with in situ measured meteorological variables across multiple biomes in Brazil. The products were evaluated against 11 flux tower stations for seven meteorological variables (air temperature, wind speed, pressure, downward shortwave and longwave radiation, and specific humidity). Stations are spread over a variety of different Brazilian climates, and daily and monthly observational averages (or totals) were compared against the gridded products. The MSE scaled to the RMSE was calculated and intercompared among different products using a ranking system. Three additional statistical metrics (the correlation contribution, the variance contribution, and the bias contribution) were also computed to provide further insight into the cause 350 of error.

5.1 Recommendations for overall product

355 It was found that ERA5-Land performs best overall for representing multiple meteorological variables at both daily and monthly scales. This finding is consistent with studies such as Decker et al. (2012), and Wang and Zeng (2012) which indicated that ERA-Interim (ERA5-Land's predecessor) generally outperformed other datasets when validated against 33 North American flux towers and 63 China Meteorological Administration (CMA) weather observation stations over the Tibetan Plateau, respectively. Similar results were reported by Jiang et al. (2020), Pelosi et al. (2020) and Zandler et al. (2020), who 360 noted that ERA5's advanced spatial and temporal resolution contributed to superior representation of meteorological conditions. However, this study confirms that no single dataset consistently outperforms others across all variables or time scales, aligning with the conclusions of Decker et al. (2012) and Wang and Zeng (2012). Therefore, the importance of regional validation on global products is underscored.

5.2 Recommendations for each variable

365 At the daily scale, ERA5-Land was found to be the most accurate for all variables except precipitation, where MSWEPv2.2 aligned more closely with observations. This finding aligns with multiple other studies which demonstrated that MSWEP exhibited strong precipitation representation, particularly in data-scarce regions (Alijanian et al. 2019; Xu et al. 2019), including South America (Moreira et al. 2019). At the monthly scale, our analysis shows that ERA5-Land best represents pressure, air temperature, longwave radiation, and specific humidity, while BNMD performs best for wind speed and 370 precipitation. Comparisons can be drawn between this study and Decker et al. (2012), who despite comparing previous versions of some of the products analysed here (ERA-Interim, ERA-40 and GLDAS1.0), they concluded ERA-Interim (the predecessor of ERA5/ERA5-Land) outperformed the other products across most variables. As mentioned above, ERA-Interim also performed well when compared to CMA measurements across the Tibetan Plateau (Wang and Zeng, 2012). The dataset achieved the best performance at both daily and monthly air temperatures whilst also demonstrating low biases and high



375 correlations in other variables, such as precipitation. Given that ERA-Interim is the predecessor of ERA5-Land, with many parallel techniques used in dataset production, similar performance can be expected in this study.
To provide a clearer understanding of the absolute deviations between observations and best-performing product, a time series comparison for air temperature and windspeed at site PDG are presented in Fig. 7. This example offers a visual interpretation of a small section of Fig. 3 while addressing the need for absolute metric comparisons. The time series reveals that ERA5-
380 Land consistently underestimates wind speed consistently by approximately 0.5 m/s, while showing only minor deviations in air temperature. Such performance is reasonable for a gridded product, highlighting its strength in capturing seasonal trends.

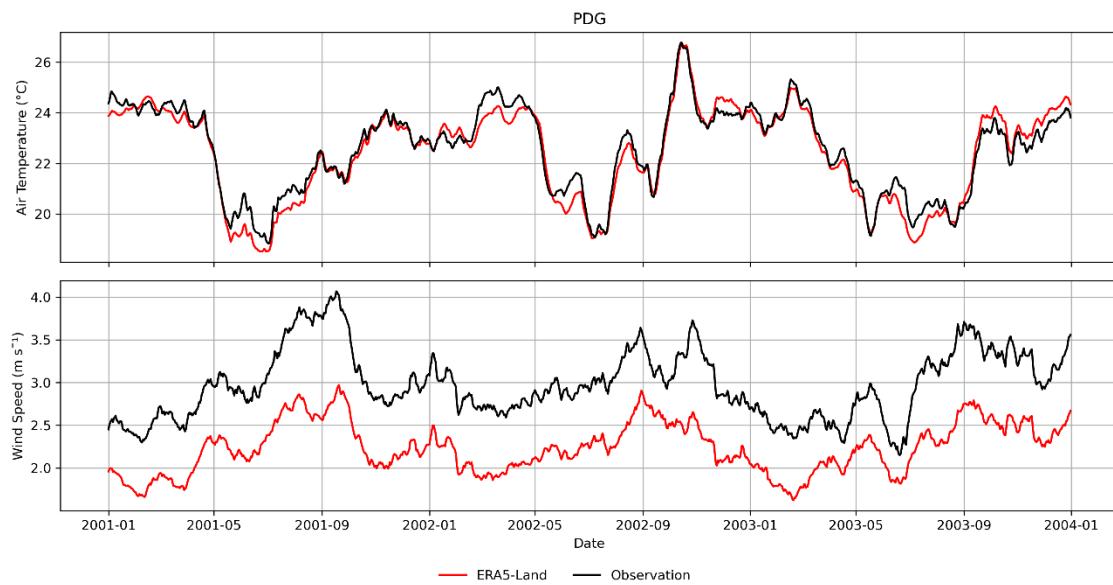


Figure 7 The 30-day rolling average for air temperature and wind speed at site PDG for observation measurements (black) and ERA5-Land (red)

Surprisingly, however, GLDAS (version 1) performed best in both daily and monthly precipitation in Wang and Zeng's 2012 study, while its successor, GLDAS2.0, ranked worst at both time scales in this analysis. This discrepancy implies that different datasets may not perform consistently on a global scale and may exhibit superiority over others depending on regional and
385 climatic contexts.

The outperformance of BNMD over MSWEPv2.2 when the temporal resolution becomes coarser, suggests that despite capturing daily patterns less well, BNMD captures the overall seasonality better than MSWEPv2.2. One explanation could be that BNMD has a greater correlation error but lower errors in variation and bias at the daily scale. When both datasets are "smoothed out" with the decrease in temporal resolution, BNMD's correlation error drops the overall MSE more than MSWEP.
390 Both products use an extensive network of rain gauges to create the gridded product, but use different methods of interpolation (as well as inevitably a few different sources). It is with interest that they outperform each other at different scales as this proves that different approaches to the creation of data products could prove suitable depending on the concerning temporal scale.



Meanwhile, GLDAS2.0 outperformed other datasets only for downward shortwave radiation at the monthly scale, consistent
395 with results from Decker et al. (2012) and Wang and Zeng (2012), where GLDAS (version 1) excelled in solar radiation on other continents. The GLDAS2.0 product is forced using the Global Meteorological Forcing Dataset from Princeton University, which combines incoming shortwave radiation from NCEP reanalysis (Kalnay et al., 1996) and NASA Langley surface radiation budget data (Cox et al., 2017), using monthly data. This could explain GLDAS2.0's superiority at the monthly scale, as the monthly signal from the forcing datasets is effectively conserved.

400 **5.3 Dominant types of error**

Our analysis reveals that the dominant sources of error vary significantly depending on the variable and time scale, reflecting the complexity of accurately capturing different meteorological factors. At finer time scales (e.g., daily), correlation error emerges as the largest contribution for most variables, including precipitation, air temperature, solar radiation, and thermal radiation. This is likely because these variables fluctuate quickly over time, making it challenging for models to maintain
405 alignment with observed temporal patterns. Conversely, errors associated with bias contribute the most to pressure and specific humidity, likely due to the stable nature of these variables and the potential accumulation of systematic offsets. Wind speed displays a more balanced distribution of error contributions from both correlation and bias, possibly reflecting the variable's high sensitivity to local topographic and atmospheric conditions, which can vary across sites. This finding mirrored observations by Decker et al. (2012) who acknowledged that the correlation contribution was more prominent at the daily and
410 sub-daily scale.

When shifting to a coarser time scale (e.g., monthly), the dominant error contributions shift as well. Specifically, the relative contribution of bias increases across all variables, while correlation contributions decrease aligning with Decker et al. (2012) who also noticed this shift. This is anticipated, as temporal averaging at a monthly scale reduces the impact of time lags inherent in correlation errors, effectively smoothing out short-term discrepancies and highlighting systematic biases. Such
415 changes underscore the importance of temporal resolution in model evaluations, as daily errors may understate or overstate the importance of correlation and bias depending on the analysis period.

For instance, the bias associated with wind speed might stem from assumptions made during vertical interpolation across datasets. In particular, assumptions about atmospheric stability or wind profile shape could introduce systematic errors that manifest as bias, especially at coarser time scales. Pressure, being relatively stable, may not capture short-term fluctuations as
420 errors but instead reveal a tendency for overprediction or underprediction that surfaces as a steady bias. Specific humidity, which depends on pressure, is similarly prone to bias errors due to its sensitivity to any pressure-related inaccuracies. These findings highlight the need to consider both temporal scale and variable characteristics in future model development and error correction approaches. Such detailed breakdowns offer a clearer understanding of the nuances in model performance, which could guide targeted improvements for specific variables and time scales.



425 **5.4 Variation in error by location and seasonality**

Our results demonstrate a clear seasonal component in the errors for precipitation and shortwave radiation, with lower errors during dry seasons and heightened errors during wetter season. Precipitation error is expected as there is higher chance for error with more rain. Solar radiation error could be associated with increased cloud cover as it follows a similar pattern and are more difficult to replicate in modelled systems such as LSMs.

430 Speculating on the latitudinal impact on error proves challenging, as no clear patterns emerge. This does not imply that latitude lacks influence; rather, other factors, such as the dominant vegetation type, may obscure potential trends. Notably, the correlation contribution in air temperature does appear to be affected, with errors generally increasing with distance from the equator.

5.5 Methodological and instrument limitations

435 While our analysis incorporated quality-controlled observational data, inherent limitations in flux tower measurements, such as instrument errors and episodic operation, remain a concern. As highlighted by Hollinger and Richardson (2005), flux tower data can deteriorate over time, which may introduce discrepancies when comparing these observations to gridded products, and in this study all data that was inside the scope of the variability, and therefore could not be rejected, was kept.

440 Additionally, reanalysis and interpolation methods differ among datasets, introducing unique biases. For instance, MSWEP and BNMD utilise distinct approaches to rain gauge data interpolation, resulting in varied precipitation accuracy depending on the region and scale.

BNMD dataset is based on the spatial interpolation of a network of meteorological stations, generally installed at a height of 1.5 m in a standard WMO grass-covered area (Xavier et al., 2016). Vertical interpolation of air temperature was not undertaken due to the complexity of modelling sub-canopy temperature gradients and the absence of sufficient high-resolution vertical 445 profile data. As a result, temperature discrepancies may have arisen, particularly with BNMD, where air temperatures at ground level were likely higher than those at forest canopy height, contributing to the observed overestimation.

Moreover, the spatial resolution of datasets appears to limit accuracy at daily time steps, where smaller-scale variability is more critical. However, this limitation becomes less prominent at monthly scales, as shown by Decker et al. (2012) and Wang and Zeng (2012), supporting our finding that temporal resolution can mitigate some spatial resolution discrepancies.

450 **6. Implications and Conclusions**

This study evaluated five high-resolution meteorological global data products against 11 flux tower observations across Brazil revealing that no single data product consistently performs best across all variables and time scales. However, higher spatial and temporal resolution products (ERA5-Land and MSWEP) generally outperform the lower resolution counterparts (GLDAS2.0, GLDAS2.1 and BNMD) at the daily scale. As an overall product, the ERA5-Land dataset outperformed the others 455 at both daily and monthly time-steps.



Decomposition of the MSE provided critical insights into the primary sources of error for each variables, underlining correlation error as the most significant contributor for variables with high temporal variability, such as air temperature and precipitation, especially at finer temporal resolutions. This decomposition analysis is instrumental in guiding data product selection for model applications, as it reveals how error sources shift with variable and temporal scale, helping users weigh the
460 importance of bias, variability and correlation error depending on the application goals.

Spanning multiple climatic zones with high-quality observational data across varied time periods, this study offers valuable insights into the robustness and applicability of each data product. The findings support the use of high-resolution reanalysis products, such as ERA5-Land and MSWEP, to enhance model predictive power; however, site-specific validation remains essential for optimal performance before dataset selection. In the absence of observational data or when time constraints limit
465 validation efforts, studies like this that validate gridded datasets across diverse climatic regions become critical.

Moreover, the results emphasise the need for careful consideration of dataset characteristics and application context when selecting a gridded data product. For instance, in applications like evapotranspiration modelling for agriculture, datasets that perform well in the dry season may be preferable (Blankenau et al., 2020). Conversely, for studies assessing long-term ecosystem responses, data products that exhibit stable performance over extended periods may be more suitable (Schymanski
470 et al. 2015). Bias correction methods and data processing steps not covered in this study may further influence dataset performance, suggesting avenues for future research. This study, alongside others, highlights that cautious, context-specific dataset selection is essential for reliable applications in environmental and climate modelling.

475

480

485



Data availability

Flux tower observation were provided by the four institutes that took part in the study and data for the early period (2000-2012) was part of the LBA-DIMP (<https://www.climatemodeling.org/lba-mip/>), while more recent data can be found as supporting material to Melo et al.'s study (2021). Four, gridded products are open access, BNMD Xavier et al. (2016), 490 (https://utexas.app.box.com/v/Xavier-etal-IJOC-DATA), GLDAS2.0 and GLDAS2.1, Rodell et al. (2004), (https://disc.gsfc.nasa.gov/datasets?keywords=GLDAS), ERA5-Land, Muñoz Sabater, J. (2019), (https://cds.climate.copernicus.eu/datasets/reanalysis-era5-land?tab=download). The precipitation dataset MSWEPv2.2 was kindly provided by Hylke Beck over personal contact (Beck et al., 2021).

Author contribution

495 JB conceptualised the study, conducted the analysis, and wrote the manuscript. Supervision and guidance were provided by RR and RW. Flux tower data across Brazil were managed and supplied by HR and DR. All authors contributed to the manuscript review and editing.

Competing Interest

The authors declare that they have no conflict of interest.

Acknowledgements

The author acknowledges funding to carry out the research from the following funding bodies; The Engineering and Physical Sciences Research Council (EPSRC) Water Informatics: Science and Engineering Centre for Doctoral Training (WISE-CDT; grant no. EP/L016214/1), the Brazilian Experimental datasets for Multi-Scale interactions in the critical zone under Extreme Drought (BEMUSED; grant no. NE/R004897/1), and partial support from the COSMIC-SWAMP project, jointly funded by 505 NERC under grant NE/W004364/1, and the São Paulo Research Foundation (FAPESP) under grant 2021/03032-7. HR acknowledges financial support, in part of FAPESP 2021/11762-5 and CTG-IAG P&D/Aneel - Sistema inteligente para benefício da qualidade das informações climáticas e ampliação da rede de estações meteorológicas no setor de energia hidro-eólica-solar do Brasil.

For the purpose of open access, the author(s) has applied a Creative Commons Attribution (CC BY) licence to any 510 Author Accepted Manuscript version arising from this submission.



References

Alijanian, M., G.R. Rakhshandehroo, A. Mishra, & Dehghani, M.: Evaluation of remotely sensed precipitation estimates using PERSIANN-CDR and MSWEP for spatio-temporal drought assessment over Iran. *Journal of Hydrology*, 579, 124189, doi:10.1016/j.jhydrol.2019.124189, 2019.

515 Allen, Richard & Pereira, L. & Raes, D. & Smith, M.: FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations. 56. 26-40, 1998.

Alvares, C. A., Stape, J. L., Sentelhas, P. C., De Moraes Gonçalves, J. L., & Sparovek, G.: Köppen's climate classification map for Brazil. *Meteorologische Zeitschrift*, 22(6), 711–728, doi:10.1127/0941-2948/2013/0507, 2013.

520 Andreae, M. O., Artaxo, P., Brandão, C., Carswell, F. E., Ciccioli, P., da Costa, A. L., Culf, A. D., Esteves, J. L., Gash, J. H. C., Grace, J., Kabat, P., Lelieveld, J., Malhi, Y., Manzi, A. O., Meixner, F. X., Nobre, A. D., Nobre, C., Ruivo, M. d. L. P., Silva-Dias, M. A. Stefani, P., Valentini, R., von Jouanne, J., Waterloo, M. J.: Biogeochemical cycling of carbon, water, energy, trace gases, and aerosols in Amazonia: The LBA-EUSTACH experiments. *Journal of Geophysical Research Atmospheres*, 107(20), doi:10.1029/2001JD000524, 2002.

Araújo, A. C.: Comparative measurements of carbon dioxide fluxes from two nearby towers in a central Amazonian rainforest: 525 The Manaus LBA site. *Journal of Geophysical Research*, 107(D20), doi:10.1029/2001jd000676, 2002

Baez-Villanueva, O.M., Zambrano-Bigiarini, M., Ribbe, L., Nauditt, A., Giraldo-Osorio, J. D., & Thinh, N.X.: Temporal and spatial evaluation of satellite rainfall estimates over different regions in Latin-America. *Atmospheric Research*, 213, 34-50, doi:10.1016/j.atmosres.2018.05.011, 2018.

530 Baldocchi, D., Falge, E., Gu, L., Olson, R., Hollinger, D., Running, S., Anthoni, P., Bermhofer, C., Davis, K., Evans, R., Fuentes, J., Goldstein, A., Katul, G., Law, B., Lee, X., Malhi, Y., Meyers, T., Munger, W., Oechel, W., Paw, K. T., Pilegaard, K., Schmid, H. P., Valentini, R., Verma, S., Vesala, T., Wilson, K., Wofsy, S.: Fluxnet: A new tool to study the temporal and spatial variability of ecosystem-scale carbon dioxide, water vapor, and energy flux densities. *Bull. Amer. Meteor. Soc.*, 82, 2415–2434, doi:10.1175/1520-0477(2001)082<2415:FANTTS>2.3.CO;2, 2001.

Beck, H. E., Wood, E. F., Pan, M. & Fisher, C. K.: MSWEP V2 Global 3-Hourly 0.1° Precipitation: Methodology and 535 Quantitative Assessment. *Bulletin of the American Meteorological Society*, 100, 473-500, doi:10.1175/BAMS-D-17-0138.1, 2019.

Beck, H. E., Pan, M., Miralles, D. G., Reichle, R. H., Dorigo, W. A., Hahn, S., Sheffield, J., Karthikeyan, L., Balsamo, G., Parinussa, R. M., van Dijk, A. I. J. M., Du, J., Kimball, J. S., Vergopolan, N., & Wood, E. F.: Evaluation of 18 satellite-and model-based soil moisture products using in situ measurements from 826 sensors. *Hydrol. Earth Syst. Sci.* 25(1), 17-40, 540 doi:10.5194/hess-25-17-2021, 2021.

Beven, K. and Westerberg, I.: On red herrings and real herrings: disinformation and information in hydrological inference. *Hydrological Processes*, 25(10), 1676-1680, doi:10.1002/hyp.7963, 2011.



Blankenau, P. A., Kilic, A., & Allen, R.: An evaluation of gridded weather data sets for the purpose of estimating reference evapotranspiration in the United States. *Agricultural Water Management*, 242, doi:10.1016/j.agwat.2020.106376, 2020.

545 Bolton, D.: The Computation of Equivalent Potential Temperature. *Monthly Weather Review*, 108, 1046–1053, doi:10.1175/1520-0493(1980)108<1046:TCOEPT>2.0.CO;2, 1980.

Bonal, D., Burban, B., Stahl, C., Wagner, F., & Héault, B.: The response of tropical rainforests to drought—lessons from recent research and future prospects. *Annals of Forest Science*, 73(1), 27–44, doi:10.1007/s13595-015-0522-5, 2016.

550 Brunke, M. A., Fairall, C. W., Zend, X., Eymard, L., and Curry, J. A.: Which Bulk Aerodynamic Algorithms are Least Problematic in Computing Ocean Surface Turbulent Fluxes? *Journal of Climate*, 16, 4, 619–635, doi:10.1175/1520-0442(2003)016<0619:WBAAAL>2.0.CO;2, 2003.

Borma, L. D. S., Da Rocha, H. R., Cabral, O. M., Von Randow, C., Collicchio, E., Kurzatkowski, D., Brugger, P. J., Freitas, H., Tannus, R., Oliveira, L. and Rennó, C. D.: Atmosphere and hydrological controls of the evapotranspiration over a floodplain forest in the Bananal Island region, Amazonia. *Journal of Geophysical Research: Biogeosciences*, 114(G1), 555 doi:10.1029/2007JG000641, 2009.

Cabral, O.M., da Rocha, H.R., Gash, J.H., Ligo, M.A., Freitas, H.C. and Tatsch, J.D.: The energy and water balance of a Eucalyptus plantation in southeast Brazil. *Journal of Hydrology*, 388(3-4), 208–216, doi:10.1016/j.jhydrol.2010.04.041, 2010.

Cabral, O. M. R., Rocha, H. R., Gash, J. H., Ligo, M. A. V., Ramos, N. P., Packer, A. P., & Batista, E. R.: Fluxes of CO₂ above a sugarcane plantation in Brazil. *Agricultural and Forest Meteorology*, 182–183, 54–66, 560 doi:10.1016/j.agrformet.2013.08.004, 2013.

Coe, M.T., Macedo, M.N., Brando, P.M., Lefebvre, P., Panday, P., Silvério, D.: The Hydrology and Energy Balance of the Amazon Basin. In: Nagy, L., Forsberg, B., Artaxo, P. (eds) *Interactions Between Biosphere, Atmosphere and Human Land Use in the Amazon Basin*. Ecological Studies, vol 227. Springer, Berlin, Heidelberg, 2016.

Cox, S.J., Stackhouse, P.W., Gupta, S.K., Mikovitz, C.J., Zhang, T.: NASA/GEWEX shortwave surface radiation budget: 565 Integrated data product with reprocessed radiance, cloud, and meteorology inputs, and new surface albedo treatment. *AIP Conference Proceedings* 1810: 090001. <https://doi.org/10.1063/1.4975541>, 2017.

De Gonçalves, L.G.G., Borak, J.S., Costa, M.H., Saleska, S.R., Baker, I., Restrepo-Coupe, N., Muza, M.N., Poulter, B., Verbeeck, H., Fisher, J.B. and Arain, M.A.: Overview of the large-scale biosphere–atmosphere experiment in Amazonia Data Model Intercomparison Project (LBA-DMIP). *Agricultural and Forest meteorology*, 182, 111–127, 570 <https://doi.org/10.1016/j.agrformet.2013.04.030>, 2013.

Decker, M., M. Brunke, Z. Wang, K. Sakaguchi, X. Zeng, & M. Bosilovich.: Evaluation of the reanalysis products from GSFC, NCEP, and ECMWF using flux tower observations. *Journal of Climate*, 25, 1916–1944, doi:10.1175/JCLI-D-11-00004.1, 2012.

Díaz, M. F., Bigelow, S., & Armesto, J. J.: Alteration of the hydrologic cycle due to forest clearing and its consequences for 575 rainforest succession. *Forest Ecology and Management*, 244(1–3), 32–40, doi:10.1016/j.foreco.2007.03.030, 2007.



Dullaart, J. C. M., Muis, S., Bloemendaal, N., & Aerts, J. C. J. H.: Advancing global storm surge modelling using the new ERA5 climate reanalysis. *Climate Dynamics*, 54(1–2), 1007–1021, doi:10.1007/s00382-019-05044-0, 2020.

Filho, A. J. P., Vemado, F., Vemado, G., Vieira Reis, F. A. G., do Carmo Giordano, L., Cerri, R. I., dos Santos, C. C., Lopes, E. S. S., Gramani, M. F., Ogura, A. T., Zaine, J. E., da Silva Cerri, L. E., Filho, O. A., D’Affonseca, F. M., & dos Santos 580 Amaral, C.: A step towards integrating CMORPH precipitation estimation with rain gauge measurements. *Advances in Meteorology*, ID 2095304, 1–24, doi:10.1155/2018/2095304, 2018.

Gampe, D.; Ludwig, R.: Evaluation of Gridded Precipitation Data Products for Hydrological Applications in Complex Topography. *Hydrology*, 4, 53, doi:10.3390/hydrology4040053, 2017.

Goulden, M. L., Miller, S. D., Da Rocha, H. R., Menton, M. C., de Freitas, H. C., e Silva Figueira, A. M., & de Sousa, C. A. 585 D.: Diel and seasonal patterns of tropical forest CO₂ exchange. *Ecological Applications*, 14(4), 42–54, doi:10.1890/02-6008, 2004.

Gupta, H.V., H. Kling, K.K. Yilmaz, & Martinez, G.F.: Decomposition of the mean squared error and NSE performance criteria: Implications for improving hydrological modelling. *Journal of Hydrology*, 377, 80–91, doi:10.1016/j.jhydrol.2009.08.003, 2009.

Hollinger, D.Y., & Richardson, A.D.: Uncertainty in eddy covariance measurements and its application to physiological models. *Tree Physiology*, 25:7, 873–885, doi:10.1093/treephys/25.7.873, 2005.

Hughes, D. A., A. Slaughter.: Daily disaggregation of simulated monthly flows using different rainfall datasets in southern Africa, *Journal of Hydrology: Regional Studies*, 4, 153–171, doi:10.1016/j.ejrh.2015.05.011, 2015.

IPCC, 2012 – Field, C.B., V. Barros, T.F. Stocker, D. Qin, D.J. Dokken, K.L. Ebi, M.D. Mastrandrea, K.J. Mach, G.-K. 595 Plattner, S.K. Allen, M. Tignor, and P.M. Midgley (Eds.) Available from Cambridge University Press, The Edinburgh Building, Shaftesbury Road, Cambridge CB2 8RU ENGLAND, 582 pp. Available from June 2012.

Iwema, J., Rosolem, R., Rahman, M., Blyth, E., and Wagener, T.: Land surface model performance using cosmic-ray and point-scale soil moisture measurements for calibration, *Hydrol. Earth Syst. Sci.*, 21, 2843–2861, <https://doi.org/10.5194/hess-21-2843-2017>, 2017.

Jiang, H., Y. Yang, Y. Bai, & Wang, H.: Evaluation of the Total, Direct, and Diffuse Solar Radiations from the ERA5 600 reanalysis data in China. *Geoscience and remote sensing letters*, 17(1), 47–51, doi:10.1109/LGRS.2019.2916410, 2020.

Jung, M., Nelson, J., Migliavacca, M., El-Madany, T., Papale, D., Reichstein, M., Walther, S., and Wutzler, T.: Technical note: Flagging inconsistencies in flux tower data, *Biogeosciences*, 21, 1827–1846, <https://doi.org/10.5194/bg-21-1827-2024>, 2024.

Kauffeldt, A., Halldin, S., Rodhe, A., Xu, C.Y. and Westerberg, I.K.: Disinformative data in large-scale hydrological 605 modelling. *Hydrol. Earth Syst. Sci.*, 17(7), 2845–2857, doi:10.5194/hess-17-2845-2013, 2013.

Liebmann, B., & D. Allured: Daily precipitation grids for South America. *Bulletin for the American Meteorological Society*, 86, 1567–1570, doi:10.1175/BAMS-86-11-1567, 2005.



Kalnay, E., Kanamitsu, M., Kistler, R., Collins, W., Deaven, D., Gandin, L., Iredell, M., Saha, S., White, G., Woollen, J., Zhu, 610 Y., Leetmaa, A., Reynolds, R., Chelliah, M., Ebisuzaki, W., Higgins, W., Janowiak, J., Mo, K. C., Ropelewski, C., Wang, J., Jenne, R., & Joseph, D.: The NCEP/NCAR 40-year reanalysis project. *Bulletin of the American Meteorological Society*, 77(3), 437–471. [https://doi.org/10.1175/1520-0477\(1996\)077<0437:TNYRP>2.0.CO;2](https://doi.org/10.1175/1520-0477(1996)077<0437:TNYRP>2.0.CO;2), 1996.

Kirkman, G. A., Gut, A., Ammann, C., Gatti, L. V., Cordova, A. M., Moura, M. A. L., Andreae, M. O., & Meixner, F. X.: Surface exchange of nitric oxide, nitrogen dioxide, and ozone at a cattle pasture in Rondônia, Brazil. *Journal of Geophysical Research*, 107(D20), doi:10.1029/2001JD000523, 2002.

Melo, D. C. D., Anache, J. A. A., Borges, V. P., Miralles, D. G., Martens, B., Fisher, J. B., et al.: Are remote sensing evapotranspiration models reliable across South American ecoregions? *Water Resources Research*, 57, e2020WR028752. <https://doi.org/10.1029/2020WR028752>, 2021.

Miralles, D. G., Gentine, P., Seneviratne, S. I., and Teuling, A. J.: Land-atmospheric feedbacks during droughts and heatwaves: 620 state of the science and current challenges, *Annals of the New York Academy of Sciences*, 1436(1), 19–35, doi:10.1111/nyas.13912, 2019.

Moreira, A.A., Ruhoff, A.L., Roberti, D.R., Souza, V.d.A., da Rocha, H. R., & de Pavia, R.C.D.: Assessment of terrestrial water balance using remote sensing data in South America. *Journal of Hydrology*, 575, 131-147, doi:10.1016/j.jhydrol.2019.05.021, 2019.

625 Moreira, V. S., Candido, L. A., Roberti, D. R., Webler, G., Diaz, M. B., de Gonçalves, L. G. G., Pousa, R., & Degrazia, G. A.: Influence of soil properties in different management systems: Estimating soybean water changes in the agro-IBIS model. *Earth Interactions*, 22(4), 1–19, doi:10.1175/EI-D-16-0033.1, 2018.

Muñoz-Sabater, J., (2019): ERA5-Land hourly data from 1981 to present. Copernicus Climate Change Service (C3S) Climate Data Store (CDS). (Accessed on 17-05-2021), 10.24381/cds.e2161bac

630 Oliveira, R. S., Bezerra, L., Davidson, E. A., Pinto, F., Klink, C. A., Nepstad, D. C., & Moreira, A.: Deep root function in soil water dynamics in cerrado savannas of central Brazil. *Functional Ecology*, 19(4), 574–581, doi:10.1111/j.1365-2435.2005.01003.x, 2005.

Pelosi, A., Terribile, F., D'Urso, G., & Chirico, G. B.: Comparison of ERA5-Land and UERRA MESCAN-SURFEX reanalysis Data with spatially interpolated weather observations for the regional assessment of reference evapotranspiration. *Water*, 12(6), 635 1669, doi:10.3390/w12061669, 2020.

Rice, A. H., Pyle, E. H., Saleska, S. R., Hutyra, L., Palace, M., Keller, M., De Camargo, P. B., Portilho, K., Marques, D. F., & Wofsy, S. C.: Carbon balance and vegetation dynamics in an old-growth amazonian forest. *Ecological Applications*, 14(4), S55-S71, doi:10.1890/02-6006, 2004.

da Rocha, H. R., Freitas, H. C., Rosolem, R., Juárez, R. I. N., Tannus, R. N., 22, M., Ligo, A., Cabral, O. M. R., & Dias, M. A. F. S.: Measurements of CO exchange over a woodland savanna 2 (Cerrado Sensu stricto) in southeast Brasil. *Biota Neotropica*, 2(1), 1–11, doi:10.1590/S1676-06032002000100009, 2002.



Rodell, M., Houser, P. R., Jambor, U., Gottschalck, J., Mitchell, K., Meng, C.-J., Arsenault, K., Cosgrove, B., Radakovich, J., Bosilovich, M., Entin, J. K., Walker, J. P., Lohmann, D., & Toll, D.: The Global Land Data Assimilation System. *Bulletin of the American Meteorological Society*, 85, 381-394, doi:10.1175/BAMS-85-3-381, 2004.

645 Sakai, R. K., Fitzjarrald, D. R., Moraes, O. L. L., Staebler, R. M., Acevedo, O. C., Czikowsky, M. J., da Silva, R., Brait, E., & Miranda, V.: Land-use change effects on local energy, water, and carbon balances in an Amazonian agricultural field. *Global Change Biology*, 10(5), 895–907, doi:10.1111/j.1529-8817.2003.00773.x, 2004.

Schymanski, S. J., Roderick, M. L., & Sivapalan, M.: Using an optimality model to understand medium and long-term responses of vegetation water use to elevated atmospheric CO₂ concentrations. *AoB Plants*, 7, plv060, 650 doi:10.1093/aobpla/plv060, 2015.

Sikder, M. S., David, C. H., Allen, G. H., Qiao, X., Nelson, E. J., & Matin, M. A.: Evaluation of Available Global Runoff Datasets Through a River Model in Support of Transboundary Water Management in South and Southeast Asia. *Frontiers in Environmental Science*, 7, doi:10.3389/fenvs.2019.00171, 2019

Sheffield, J., G. Goteti, & Wood, E.F.: Development of a 50-Year High-Resolution Global Dataset of Meteorological Forcings for Land Surface Modeling. *Journal of Climate*, 19, 3088-3111, doi:10.1175/JCLI3790.1, 2006.

Shuttleworth, W.J.: *Terrestrial hydrometeorology*. John Wiley & Sons, 2012.

Soti, V., Puech, C., Lo Seen, D., Bertran, A., Vignolles, C., Mondet, B., Dessay, N., & Tran, A.: The potential for remote sensing and hydrologic modelling to assess the spatio-temporal dynamics of ponds in the Ferlo Region (Senegal). *Hydrol. Earth Syst. Sci.*, 14(8), 1449–1464, doi:10.5194/hess-14-1449-2010, 2010.

660 Syed, T. H., Famiglietti, J. S., Rodell, M., Chen, J., & Wilson, C. R.: Analysis of terrestrial water storage changes from GRACE and GLDAS. *Water Resources Research*, 44, W02433, doi:10.1029/2006WR005779, 2008.

Terzago, S., Andreoli, V., Arduini, G., Balsamo, G., Campo, L., Cassardo, C., Cremonese, E., Dolia, D., Gabellani, S., von Hardenberg, J., Morra Di Cella, U., Palazzi, E., Piazzesi, G., Pogliotti, P., & Provenzale, A.: Sensitivity of snow models to the accuracy of meteorological forcings in mountain environments. *Hydrol. Earth Syst. Sci.*, 24(8), 4061–4090, doi:10.5194/hess-665 24-4061-2020, 2020.

Thornton, P. E., Shrestha, R., Thornton, M., Kao, S. C., Wei, Y., & Wilson, B. E.: Gridded daily weather data for North America with comprehensive uncertainty quantification. *Scientific Data*, 8, 190, doi: 10.1038/s41597-021-00973-0, 2021.

Tian, S., Renzullo, L. J., van Dijk, A. I. J. M., Tregoning, P., & Walker, J. P. Global joint assimilation of GRACE and SMOS for improved estimation of root-zone soil moisture and vegetation response. *Hydrol. Earth Syst. Sci.*, 23, 1067–1081, 670 doi:10.5194/hess-23-1067-2019, 2019.

Viana, F. de S., J., Maria Gico Lima Montenegro, S., Barbosa da Silva, B., Srinivasan, R., Augusto Guimarães Santos, C., Cezar dos Santos Araújo, D., & Gadelha Tavares, C.: Evaluation of gridded meteorological datasets and their potential hydrological application to a humid area with scarce data for Pirapama River basin, northeastern Brazil, *Theoretical and Applied Climatology*, 145, 393-419, doi:10.1007/s00704-021-03628-7 , 2021.



675 Vissa, N. K., Anandh, P. C., Behera, M. M., & Mishra, S.: ENSO-induced groundwater changes in India derived from GRACE and GLDAS. *Journal of Earth System Science*, 128(115), 1-9, doi:10.1007/s12040-019-1148-z, 2019.

Wagner, S., Widmann, M., Jones, J. et al.: Transient simulations, empirical reconstructions and forcing mechanisms for the Mid-holocene hydrological climate in southern Patagonia. *Clim Dyn* 29, 333–355, doi:10.1007/s00382-007-0229-x.

Wang A. & Zeng X.: Evaluation of multireanalysis products with in situ observations over the Tibetan Plateau. *Journal of Geophysical Research*, 117:D05102, doi:10.1029/2011JD016553, 2012.

680 Webler, G., Roberti, D. R., Cuadra, S. V., Moreira, V. S., & Costa, M. H.: Evaluation of a dynamic agroecosystem model (Agro-IBIS) for soybean in Southern Brazil. *Earth Interactions*, 16(12), 1–15. doi:10.1175/2012EI000452.1, 2012.

Xavier, A. C., King, C. W. & Scanlon B. R.: Daily gridded meteorological variables in Brazil (1980–2013). *International Journal of Climatology*, 36, 2644-2659, doi:10.1002/joc.4518, 2016.

685 Xi, Y., S. Peng, P. Ciais, and Y. Chen.: Future impacts of climate change on inland Ramsar wetlands. *Nat. Clim. Change*, 11, 45–51, doi:10.1038/s41558-020-00942-2, 2021.

Xu, Z., Wu, Z., He, H., Wu, X., Zhou, J., Zhang, Y., Guo, X.: Evaluating the accuracy of MSWEP V2.1 and its performance for drought monitoring over mainland China. *Atmospheric Research*, 226, 17-31, doi:10.1016/j.atmosres.2019.04.008, 2019.

Zandler, H., Senftl, T., & Vanselow K.A.: Reanalysis datasets outperform other gridded climate products in vegetation change 690 analysis in peripheral conservation areas of central Asia. *Nature: Scientific Reports*, 10:22446, doi:10.1038/s41598-020-79480-y, 2020.