



Chloric Acid-Driven Nucleation Enhanced by
Dimethylamine and Sulfuric acid in the Arctic:
Mechanistic Study
Shengming Wang ¹ , Huidi Zhang ² , Xiangli Shi ² , Qingzhu Zhang ^{1,*} ,
Wenxing Wang ¹ , Qiao Wang ¹
¹ Environment Research Institute, Shandong University,
Qingdao 266237, P. R. China
² College of Geography and Environment, Shandong Normal
University, Jinan 250014, PR China
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23 Abstract

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25 Chlorine radicals are strong oxidizing agents in the atmosphere, and the process of chlorine oxidation results in the formation of chloric acid (HClO₃, CA). Recent 26 studies have shown that CA is prevalent in the Arctic boundary layer. However, the 27 contribution of chlorine-containing species to oceanic new particle formation (NPF) 28 has not been fully revealed. It is expected that CA is involved in the oceanic 29 nucleation process. In this study, the enhancement of CA-based NPF by 30 dimethylamine (DMA) and sulfuric acid (SA) was comparatively investigated at the 31 molecular level using density-functional theory (DFT) and atmospheric cluster 32 dynamics simulation (ACDC). The results show that DMA can form clusters with CA 33 through hydrogen bonding, halogen bonding and proton transfer, which reduces the 34 energy barrier for CA-based cluster formation and significantly improves the 35 thermodynamic stability of CA clusters. The cluster formation rate of CA-DMA 36 cluster system is higher than that of the CA-SA cluster system. The CA-DMA cluster 37 system in the Arctic atmosphere contributes to NPF. These findings may help to 38 39 reveal some of the missing sources of the Arctic NPF. The present study contributes to a deeper understanding of the influence of oceanic chlorine-containing constituents 40 on the oceanic NPF. 41

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1 Introduction

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Marine aerosols as the main natural aerosol system have a major global impact 46 by regulating the radiative balance and climate of clouds(O' dowdg). New particle 47 48 formation (NPF) contributes to more than half of the global cloud condensation nuclei, which in turn contributes to cloud formation (Gordon et al., 2017; Takegawa et al., 49 50 2020; Williamson et al., 2019; Zhang et al., 2012). Compared with clouds over land, 51 ocean clouds cover a wider area and significantly increase the albedo of the ocean, so 52 ocean clouds contribute more to the climate system (Merikanto et al., 2009; Wood, 53 2012; Zheng et al., 2021). Sulfuric acid (SA, H₂SO₄), methane sulfonic acid (MSA, CH₃HSO₃), and iodic acid (IA, HIO₃) are generally considered to contribute to the 54 55 formation of oceanic particles (Hodshire et al., 2019; Yin et al., 2021; Facchini et al., 2008; Perraud et al., 2015; Arquero et al., 2017; Hopkins et al., 2008). However, there 56 is still a significant difference between the particle formation rates observed in the 57 field and those predicted by simulation (Kirkby et al., 2011; Kirkby et al., 2016; 58 59 Zhang et al., 2004; Ehn et al., 2014; Dawson et al., 2012). Therefore, it is necessary to consider whether other gaseous precursors are involved in NPF to narrow the gap 60 between experiments and simulations. 61 Compared to the main atmospheric oxidants, hydroxyl radicals, chlorine radicals 62 63 act as strong oxidants in the polar troposphere at relatively high concentration levels (Stone et al., 2012). The active chlorine cycle in the Arctic boundary layer during the 64 spring after polar sunrise depletes O₃ in the region (Custard et al., 2016; Foster et al., 65





2001; Thompson et al., 2015). Chloric acid (CA) has no photoactivity and CA is 66 67 ubiquitous in the spring in the Arctic, with concentrations estimated to range from 1×10^5 to 7×10^6 molecules cm⁻³ (Tham et al., 2023). CA was not found in the 68 particle phase of the aerosol, thus it is difficult to determine whether CA is involved in 69 70 the NPF phase. Many studies have shown that atmospheric bases such as methylamine (MA), 71 72 dimethylamine (DMA), trimethylamine (TMA) and ammonia can effectively enhance 73 SA-based NPF (Yao et al., 2018; Almeida et al., 2013). Although amines emit 10–20 74 times less than ammonia in the ocean, amines can effectively form clusters with 75 substances such as SA in the ocean (Myriokefalitakis et al., 2010; Semeniuk and Dastoor, 2018; Almeida et al., 2013). Of these amines, DMA has been found as a 76 77 component of marine secondary aerosols (Facchini et al., 2008). Widely dispersed DMA has an atmospheric concentration of 0.4 - 10 pptv over the ocean and plays a 78 key role in marine NPF(Van Pinxteren et al., 2019). DMA has been identified as the 79 strongest enhancing atmospheric amine for SA and IA-driven NPF (Olenius et al., 80 81 2017; Ning et al., 2022). Thus, DMA may have a higher enhancing potential (EP) than NH₃ for CA-based NPF. 82 Sulfuric acid (H₂SO₄, SA) has been detected in both gas and particulate phases in 83 polluted coastal areas of China(Zhu et al., 2019; Yu et al., 2019). It is noteworthy that 84 the concentration of SA is two orders of magnitude higher (up to 10⁸ molecules cm⁻³) 85 in the coastal polluted areas due to urban air pollution compared to the clean marine 86 atmosphere(Zhu et al., 2019). Considering its strong nucleation ability(Sipilä et al., 87





88 2010; Faloona, 2009), it is likely that SA molecules nucleate in marine regions along

with CA.

Using density functional theory (DFT) and the Atmospheric Clusters Dynamic 90 Code (ACDC), the involvement of DMA and SA in the initial phase of CA-based 91 92 NPF has been investigated. We obtained the minimum free energy structures of the (CA)₁₋₄(DMA)₁₋₄ and (CA)₁₋₄(SA)₁₋₄ clusters. The temperatures used in this study are 93 94 within the temperature range of the atmospheric boundary layer in the ordinary 95 range(Miřijovský and Langhammer, 2015). The concentration of CA was estimated to be in the range of $1.0 \times 10^6 - 1.0 \times 10^8$ molecules cm⁻³ based on measured data. 96 97 Further study of CA-DMA clusters under Arctic atmospheric conditions and the corresponding thermodynamic data used as input to the ACDC reveals the growth 98

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2 Computational methods

pathways and formation rates of the clusters.

Configurational Sampling

We employed a multi-step global minimum sampling scheme to search for the global minimum of $(CA)_{1-4}(DMA)_{1-4}$ and $(CA)_{1-4}(SA)_{1-4}$ clusters. In this study, the initial structures of 1000-10000 $(CA)_{1-4}(DMA)_{1-4}$ and $(CA)_{1-4}(SA)_{1-4}$ clusters were randomly generated using the ABCluster software to determine their global minima (clusters with the lowest Gibbs free energies). In the multistep sampling scheme, the geometry optimization is performed at the PM7, ω B97X-D/6-31+G(d,p) and ω B97X-D/6-31++G(d,p) levels of theory, and the single-point energy calculations are

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110 performed at the DLPNO-CCSD(T)/aug-cc-pVTZ level of theory based on the ωB97X-D/6-31++G(d,p) theory level is performed on the geometry. The GAUSSIAN 111 112 09 program package(Frisch et al., 2016) was used to perform the PM7 and ωB97X-D calculations. DLPNO-CCSD(T) calculations were performed in the ORCA 4.0.0 113 114 program(Neese, 2012). For convergence problems and failures such as ending with a false frequency in the optimization of (CA)₁₋₄(DMA)₁₋₄ and (CA)₁₋₄(SA)₁₋₄ cluster 115 116 geometries, the initial structures will be modified and re-optimized until the 117 optimization is successful. The free energy of formation (ΔG) of individual clusters is 118 calculated at different temperatures 238, 258, and 278 K. The ΔG of individual 119 clusters is calculated at different temperatures. The structures of (SA)₁₋₄ and (DMA)₁₋₄ clusters were obtained from previous studies and are recalculated here(Xie 120 121 et al., 2017).

Atmospheric Cluster Dynamics Code (ACDC) Simulation

Time-evolving cluster formation rates, steady-state concentrations, and growth paths for $(CA)_{1-4}(DMA)_{1-4}$ and $(CA)_{1-4}(SA)_{1-4}$ clusters were calculated using ACDC(Mcgrath et al., 2012). There is good agreement between the conclusions of the ACDC simulations and the experimental results obtained using the birth and death equations (Almeida et al., 2013; Lu et al., 2020; Kürten et al., 2018). In this study, ACDC simulations were performed to model the formation process of CA-DMA and CA-SA neutral clusters without considering the effects of charge and water. The $(CA)_5(DMA)_5$ clusters are set as boundary clusters (see Supporting Information (SI) for details). The concentration ranges of [CA], [SA] and [DMA] were set to 10^6-10^8





 cm^{-3} , $10^6 - 10^8$ cm⁻³ and 0.1–100 ppt, respectively.

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3 Results and discussion

3.1 Cluster Structures and Cluster Formation Free Energy

136 To evaluate the thermodynamic stability of the formed CA-DMA clusters, the formation free energies (ΔG , kcal mol⁻¹) at 278 K are calculated for the 137 138 $(CA)_{1-4}(DMA)_{1-4}$ clusters at DLPNO-CCSD(T)/aug-cc-pVTZ//ωB97X-D/6-31++G(d,p) level of theory. As shown 139 140 in Figure. 1, hydrogen bonds play an important role in the formation of CA-DMA 141 clusters. Proton transfer reactions are not observed in the pure (CA)₂ and (CA)₃ clusters, whereas spontaneous proton transfer reactions are observed in all 142 143 $(CA)_{1-4}(DMA)_{1-4}$ clusters. For most $(CA)_{1-4}(DMA)_{1-4}$ clusters, protons are transferred from CA to DMA. 144 In order to compare the thermodynamic stability of the formed CA-DMA and 145 CA-SA clusters, we calculated the formation Gibbs free energy values (ΔG , kcal 146 147 mol^{-1}) for the $(CA)_{1-4}(DMA)_{1-4}$ clusters 278 K at the DLPNO-CCSD(T)/aug-cc-pVTZ// ω B97X-D/6-31++G(d,p) level of theory. As shown 148 149 in Figure. 1, hydrogen bonding plays an important role in the formation of CA-DMA clusters. Proton transferring wasn't observed in pure (CA)2 and (CA)3 clusters, 150 151 whereas spontaneous proton transfer reactions were present in all (CA)₁₋₄(DMA)₁₋₄ clusters. For most (CA)₁₋₄(DMA)₁₋₄ clusters, protons are transferred from CA to 152 DMA, forming Cl-O...H-N hydrogen bonding, accompanied by the production of 153





154 ClO₃⁻ negative ions and DMA⁺ ions. In CA-DMA clusters containing two or more chlorine 155 atoms, including $(CA)_2(DMA)_1$ $(CA)_3(DMA)_1$ $(CA)_3(DMA)_1$ (CA)₃(DMA)₁, (CA)₃(DMA)₄, (CA)₄(DMA)₁, (CA)₄(DMA)₂, (CA)₄(DMA)₃, and 156 (CA)4(DMA)4 clusters, O-O...O-Cl halogen bonds and Cl-O...H-N hydrogen bonds 157 158 together stabilize these clusters described above. Halogen bonds are not present in clusters containing single CA atom and in (CA)₂(DMA)₂, (CA)₂(DMA)₃, 159 160 (CA)₂(DMA)₄, as well as (CA)₃(DMA)₃ clusters. For the CA-SA clusters, the ClO₃⁻ 161 negative ions generated in the CA-DMA cluster system were not found in the CA-SA 162 cluster system because proton transfer do not occur. In contrast to the O-Cl...O-Cl 163 halogen bond found in the CA-DMA cluster system, the CA-SA cluster adds S-O...Cl-O halogen bond (in the Figure S2). 164 165 The formation Gibbs free energy values of (A) (CA)₁₋₄(DMA)₁₋₄ and (B) (CA)₁₋₄(SA)₁₋₄ at 278 K and 1 atm are shown in Figure. 2. The Gibbs free energy 166 values for the formation of (CA)₂, (CA)₃, and (CA)₄ clusters in the pure acid system 167 are 1.31, 2.52, and 5.37 kcal mol⁻¹, respectively, implying that at 278 K, pure CA 168 169 clusters are thermodynamically difficult to form. The ΔG values of (CA)₁(SA)₄ and pure SA clusters are lower than the corresponding ΔG values of the corresponding 170 CA-DMA system. The difference between the ΔG values of (CA)₄(DMA)₄ cluster and 171 (CA)₄(SA)₄ cluster is the largest, up to 61.62 kcal mol⁻¹. (CA)₁(SA)₁ and 172 173 (CA)₁(DMA)₁ are both very important in their respective cluster systems, and their ΔG values are -2.62 and 10.08 kcal mol⁻¹, respectively⁻¹. The ΔG values of (CA)₁₋₄ 174 clusters are 10.08 - 28.16 kcal mol⁻¹ higher than those of the corresponding 175





176 (CA)₁₋₄(DMA)₁ clusters, suggesting that pure CA clusters may grow by collision with 177 DMA. As the size of CA-DMA clusters increases, the clusters gradually form a 178 cage-symmetric structure. The ΔG values of the majority of clusters in the CA-DMA 179 system are 3.55 – 61.62 kcal mol⁻¹ lower than the corresponding ΔG values of the 180 CA-SA system. This indicates that the CA-DMA cluster system is more thermally 181 stable compared to the CA-SA cluster system.

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3.1 Evaporation Rates and Cluster Stability

184 The total evaporation rate $(\sum \gamma, s^{-1})$ of $(CA)_{1-4}(DMA)_{1-4}$ clusters and 185 $(CA)_{1-4}(SA)_{1-4}$ clusters formed at T = 278 K is shown in Figure. 3. The smaller value of $\sum y$ means that the stability of CA-DMA clusters is higher and the clusters shrink 186 187 further. The smaller value of $\sum y$ implies the higher stability of CA-DMA clusters and further contraction of the clusters. The clusters with the same number of CA 188 molecules and the number of DMA molecules include (CA)₁(DMA)₁, (CA)₂(DMA)₂, 189 (CA)₃(DMA)₃, and (CA)₄(DMA)₄ clusters, which have the values of $\sum \gamma$ of 10^2 , 2 × 190 10^{-4} , 5×10^{-1} , and 3×10^{1} s⁻¹, respectively. (CA)₂(DMA)₂ cluster has the lowest $\Sigma \gamma$ 191 value, implying that $(CA)_2(DMA)_2$ cluster is the most stable cluster in the "4 × 4" 192 box system of CA-DMA clusters. For clusters with different numbers of CA and 193 DMA molecules, the $\sum \gamma$ value of (CA)₂(DMA)₁ cluster is significantly lower than that 194 195 of other clusters, which indicates that (CA)₂(DMA)₁ cluster has a high probability of competing for the growth path of nucleation. In this study, we compare the 196 evaporation rates of clusters from the CA-DMA system with those from the CA-SA 197

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198 system at 278K. In contrast to the other clusters, the evaporation rate of (CA)₁(SA)₄ 199 clusters is lower than that of the corresponding (CA)₁(DMA)₄ clusters. The evaporation rates of most (CA)₁₋₄(DMA)₁₋₄ clusters are much smaller than those of 200 the corresponding (CA)₁₋₄(SA)₁₋₄ clusters, indicating that the CA-DMA cluster 201 202 system is kinetically more stable than the CA-SA cluster system. 203 204 3.2 Cluster Formation Rates and Steady-State Cluster Concentrations 205 Cluster formation rate (J) and steady-state CA dimer concentration ($\sum [(CA)_2]$) 206 are important indicators for assessing the enhancement potential of DMA for 207 CA-based nucleation. Figure 4 shows the variation of $\Sigma[(CA)_2]$ and J values at 278 K with CA concentration ([CA] = $10^6 - 10^8$ cm⁻³) and DMA concentration ([DMA] = 1, 208 209 10 and 100 ppt). The J value of the CA-DMA system showed a positive correlation with CA and DMA concentrations as CA and DMA concentrations increased. The 210 dependence of the cluster formation rate on the DMA concentration does not decrease 211 with increasing CA concentration, which means that the dependence of the system on 212 213 DMA does not saturate when the CA concentration is high. The $\sum [(CA)_2]$ and J of the CA-DMA system with the full range of acid-base concentrations considered (CA: 106 214 - 10⁸ molecules cm⁻³, DMA: 1, 10 and 100 ppt) were significantly higher than the 215 CA-SA system (Figure. S4). However, at high concentrations of [CA] = 1×10^7 216

molecules cm⁻³ and [DMA] = 10 ppt, the J value of the CA-DMA cluster system only

reaches 6.70×10^{-7} cm⁻³ s⁻¹. The contribution of the CA-DMA cluster system to the

NPF is not significant under the atmospheric conditions of 278 K. In addition, the





220 CA-PA cluster system has a lower J value (Figure. S15). This may be due to the weak bond energies of Cl-O...Cl-O halogen bonds in the process of CA-PA nucleation. 221 To further systematically explore the effect of temperature on the J of the 222 CA-DMA cluster system, Figure. 5 shows the simulated J at other temperatures (238 223 and 258 K), $[CA] = 10^6 - 10^8$ molecules cm⁻³, [DMA] = 0.1 ppt (red line), 1 ppt 224 (green line), 10 ppt (blue line), and 100 ppt (black line). A comparison of the 225 226 simulations at 258 K (Figure. 5a) and 238 K (Figure. 5b) reveals that the decrease in 227 temperature further increases the J value of the CA-DMA cluster system to a higher 228 level. However, at a low temperature (258 K), the J values of the CA-DMA cluster system do not reach higher levels at high concentrations of [CA] = 1×10^7 molecules 229 cm⁻³ and [DMA] = 10 ppt. The J values of the CA-DMA cluster system were further 230 231 investigated under cold Arctic conditions (238 K). It was found that the J value of CA-DMA cluster system at 238 K atmospheric condition was significantly higher 232 than that of CA-DMA cluster system at other higher temperature conditions, which 233 was mainly due to the fact that the low temperature attenuates the evaporation of 234 235 CA-DMA clusters. 236 3.3 Cluster growth pathway 237 Figure 6 shows the growth paths of CA-DMA and CA-SA clusters at 278 K, [CA] 238 = 10^6 cm⁻³, [DMA] = 1 ppt, and [SA] = 10^6 cm⁻³. The first step in the growth of 239 CA-DMA clusters is the collision of a CA molecule and a DMA molecule to form a 240 (CA)₁(DMA)₁ cluster. There are two growth paths for (CA)₁(DMA)₁ clusters: a CA 241





242 molecule collides and combines with a (CA)₁(DMA)₁ cluster to form a (CA)₂(DMA)₁ 243 cluster, and a DMA molecule is subsequently added to form a (CA)₂(DMA)₂ cluster; a (CA)₁(DMA)₁ cluster combines with another (CA)₁(DMA)₁ cluster to form a 244 (CA)2(DMA)2 cluster. This growth pattern is mainly due to the high stability of 245 246 (CA)2(DMA)1 clusters. After the (CA)2(DMA)2 clusters, the growth route of the CA-DMA cluster system extends along the direct binding to the (CA)₁(DMA)₁ 247 248 clusters. For (CA)₃(DMA)₃ clusters, one route is the addition of acid and base, and the 249 other route is the direct binding to (CA)₁(DMA)₁ clusters to produce (CA)₄(DMA)₄ 250 clusters. 251 The growth pathway of the CA-SA cluster system differs considerably from that of the CA-DMA cluster system. The cluster initially formed in the CA-SA cluster 252 253 system is the (CA)₂ cluster not the (CA)₁(SA/DMA)₁ cluster in the CA-DMA system. 254 After the generation of $(CA)_2$ cluster, the system can generate $(CA)_3(SA)_1$ clusters by successive addition of acid molecules. Subsequently, (CA)₃(SA)₁ clusters combine 255 with (CA)₁(SA)₁ clusters to generate (CA)₃(SA)₂ clusters. The final (CA)₄(SA)₄ 256 257 cluster of the system is generated by collision of a (CA)₃(SA)₂ cluster with an (SA)₂ cluster. 258 259 3.4 Atmospheric implications and conclusions 260 261 In this paper, combination method of quantum chemistry and ACDC were used to elucidate the molecular structure mechanism of DMA and SA enhancing role of 262 CA nucleation by comparing the CA-DMA and CA-SA nucleation systems. Proton 263







transfer was observed in all (CA)₁₋₄(DMA)₁₋₄ clusters while no proton transfer 264 265 occurred within (CA)₁₋₄(SA)₁₋₄ clusters. The ClO₃⁻ groups generated by the deprotonation of CA are involved in the formation of at least one hydrogen bond. 266 Hydrogen and halogen bonds together stabilize the CA-DMA and CA-SA nucleation 267 268 systems. The vast majority of CA-SA cluster systems have higher ΔG values than the corresponding CA-DMA cluster systems. The cluster formation rates of the pure 269 270 CA-PA and CA-SA nucleation systems are relatively low, and the contribution of 271 DMA to CA nucleation is stronger than that of SA. The CA-DMA nucleation system 272 contributes to the NPF at low Arctic temperatures. Clusters with the same number of 273 CA and DMA molecules ((CA)₁(DMA)₁, (CA)₂(DMA)₂, (CA)₃(DMA)₃, and (CA)₄(DMA)₄ clusters) play a key role in the growth path of CA-DMA clusters. This 274 275 study is important for a deeper understanding of CA Arctic atmospheric nucleation. The study clarifies the role of CA in the marine NPF and reveals the mechanism 276 of DMA acting as a key enhancer to CA-based NPF through intermolecular 277 interactions. The results suggest that CA-DMA synergistic nucleation in the Arctic 278 279 atmosphere is an under-recognized source of NPF, which provides a theoretical basis for improving regional climate models (e.g., cloud condensation nucleation prediction) 280 and assessing the Arctic aerosol-climate feedback effect. The current simulations do 281 not take into account charge effects, the involvement of water molecules, and the 282 283 influence of complex atmospheric matrices (e.g., organic matter). In the future, it is necessary to validate the simulation results with field observations and extend it to 284 multi-component (e.g., IA/SA/DMA mixing) nucleation systems in order to quantify 285





286	the contribution of chlorine-containing substances to the global NPF in a more
287	comprehensive way.
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289	Data availability
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291	All data supported the paper are available from the article, supplementary information,
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308 References

- 310 Almeida, J., Schobesberger, S., Kürten, A., Ortega, I. K., Kupiainen-Määttä, O., Praplan, A. P.,
- 311 Adamov, A., Amorim, A., Bianchi, F., and Breitenlechner, M.: Molecular understanding of sulphuric
- acid-amine particle nucleation in the atmosphere, Nature, 502, 359-363, 2013.
- 313 Arquero, K. D., Xu, J., Gerber, R. B., and Finlayson-Pitts, B. J.: Particle formation and growth from
- 314 oxalic acid, methanesulfonic acid, trimethylamine and water: a combined experimental and
- theoretical study, Physical Chemistry Chemical Physics, 19, 28286-28301, 2017.
- 316 Custard, K. D., Pratt, K. A., Wang, S., and Shepson, P. B.: Constraints on Arctic atmospheric
- 317 chlorine production through measurements and simulations of Cl2 and ClO, Environ. Sci. Technol.,
- 318 50, 12394-12400, 2016.
- 319 Dawson, M. L., Varner, M. E., Perraud, V., Ezell, M. J., Gerber, R. B., and Finlayson-Pitts, B. J.:
- 320 Simplified mechanism for new particle formation from methanesulfonic acid, amines, and water
- 321 via experiments and ab initio calculations, Proceedings of the National Academy of Sciences, 109,
- 322 18719-18724, 2012.
- 323 Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F.,
- 324 Tillmann, R., and Lee, B.: A large source of low-volatility secondary organic aerosol, Nature, 506,
- 325 476-479, 2014.
- 326 Facchini, M. C., Decesari, S., Rinaldi, M., Carbone, C., Finessi, E., Mircea, M., Fuzzi, S., Moretti, F.,
- 327 Tagliavini, E., and Ceburnis, D.: Important source of marine secondary organic aerosol from
- 328 biogenic amines, Environmental science & technology, 42, 9116-9121, 2008.
- 329 Faloona, I.: Sulfur processing in the marine atmospheric boundary layer: A review and critical
- assessment of modeling uncertainties, Atmospheric Environment, 43, 2841-2854, 2009.
- 331 Foster, K. L., Plastridge, R. A., Bottenheim, J. W., Shepson, P. B., Finlayson-Pitts, B. J., and Spicer, C.
- W:: The role of Br2 and BrCl in surface ozone destruction at polar sunrise, Science, 291, 471-474,
- 333 2001.
- Frisch, M., Trucks, G., Schlegel, H., Scuseria, G., Robb, M., Cheeseman, J., Scalmani, G., Barone, V.,
- 335 Petersson, G., and Nakatsuji, H.: Gaussian 16, Revision A. 03, Gaussian, Inc., Wallingford CT, 3,
- 336 2016.
- 337 Gordon, H., Kirkby, J., Baltensperger, U., Bianchi, F., Breitenlechner, M., Curtius, J., Dias, A.,
- 338 Dommen, J., Donahue, N. M., and Dunne, E. M.: Causes and importance of new particle formation
- in the present-day and preindustrial atmospheres, Journal of Geophysical Research: Atmospheres,
- 340 122, 8739-8760, 2017.
- 341 Hodshire, A. L., Campuzano-Jost, P., Kodros, J. K., Croft, B., Nault, B. A., Schroder, J. C., Jimenez, J.
- 342 L., and Pierce, J. R.: The potential role of methanesulfonic acid (MSA) in aerosol formation and
- growth and the associated radiative forcings, Atmospheric Chemistry and Physics, 19, 3137-3160,
- 344 2019.
- 345 Hopkins, R. J., Desyaterik, Y., Tivanski, A. V., Zaveri, R. A., Berkowitz, C. M., Tyliszczak, T., Gilles, M.
- 346 K., and Laskin, A.: Chemical speciation of sulfur in marine cloud droplets and particles: Analysis of
- 347 individual particles from the marine boundary layer over the California current, Journal of
- 348 Geophysical Research: Atmospheres, 113, 2008.
- 349 Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagné, S., Ickes, L.,





- 350 and Kürten, A.: Role of sulphuric acid, ammonia and galactic cosmic rays in atmospheric aerosol
- 351 nucleation, Nature, 476, 429-433, 2011.
- 352 Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M., Simon, M.,
- 353 Yan, C., and Almeida, J.: Ion-induced nucleation of pure biogenic particles, Nature, 533, 521-526,
- 354 2016.
- Kürten, A., Li, C., Bianchi, F., Curtius, J., Dias, A., Donahue, N. M., Duplissy, J., Flagan, R. C., Hakala,
- 356 J., and Jokinen, T.: New particle formation in the sulfuric acid-dimethylamine-water system:
- 357 reevaluation of CLOUD chamber measurements and comparison to an aerosol nucleation and
- growth model, Atmospheric Chemistry and Physics, 18, 845-863, 2018.
- 359 Lu, Y., Liu, L., Ning, A., Yang, G., Liu, Y., Kurtén, T., Vehkamäki, H., Zhang, X., and Wang, L.:
- 360 Atmospheric sulfuric acid dimethylamine nucleation enhanced by trifluoroacetic acid,
- 361 Geophysical Research Letters, 47, e2019GL085627, 2020.
- 362 McGrath, M., Olenius, T., Ortega, I., Loukonen, V., Paasonen, P., Kurtén, T., Kulmala, M., and
- 363 Vehkamäki, H.: Atmospheric Cluster Dynamics Code: a flexible method for solution of the
- birth-death equations, Atmospheric Chemistry and Physics, 12, 2345-2355, 2012.
- 365 Merikanto, J., Spracklen, D., Mann, G., Pickering, S., and Carslaw, K.: Impact of nucleation on
- 366 global CCN, Atmospheric Chemistry and Physics, 9, 8601-8616, 2009.
- 367 Miřijovský, J. and Langhammer, J.: Multitemporal monitoring of the morphodynamics of a
- 368 mid-mountain stream using UAS photogrammetry, Remote sensing, 7, 8586-8609, 2015.
- 369 Myriokefalitakis, S., Vignati, E., Tsigaridis, K., Papadimas, C., Sciare, J., Mihalopoulos, N., Facchini,
- 370 M., Rinaldi, M., Dentener, F., and Ceburnis, D.: Global modeling of the oceanic source of organic
- 371 aerosols, Adv, Meteorol, 939171, 2010.
- Neese, F.: The ORCA program system, Wiley Interdisciplinary Reviews: Computational Molecular
- 373 Science, 2, 73-78, 2012.
- Ning, A., Liu, L., Zhang, S., Yu, F., Du, L., Ge, M., and Zhang, X.: The critical role of dimethylamine in
- 375 the rapid formation of iodic acid particles in marine areas, npj Climate and Atmospheric Science,
- 376 5, 92, 2022
- 377 O'DowdG, C.: de Leeuw, ", Marine aerosol production: a review of the current knowledge,
- 378 1753-1774,
- Olenius, T., Halonen, R., Kurtén, T., Henschel, H., Kupiainen-Määttä, O., Ortega, I. K., Jen, C. N.,
- 380 Vehkamäki, H., and Riipinen, I.: New particle formation from sulfuric acid and amines: Comparison
- 381 of monomethylamine, dimethylamine, and trimethylamine, Journal of Geophysical Research:
- 382 Atmospheres, 122, 7103-7118, 2017.
- Perraud, V., Horne, J. R., Martinez, A. S., Kalinowski, J., Meinardi, S., Dawson, M. L., Wingen, L. M.,
- 384 Dabdub, D., Blake, D. R., and Gerber, R. B.: The future of airborne sulfur-containing particles in the
- $absence\ of\ fossil\ fuel\ sulfur\ dioxide\ emissions,\ Proceedings\ of\ the\ National\ Academy\ of\ Sciences,$
- 386 112, 13514-13519, 2015.
- 387 Semeniuk, K. and Dastoor, A.: Current state of aerosol nucleation parameterizations for
- 388 air-quality and climate modeling, Atmos. Environ., 179, 77-106, 2018.
- 389 Sipilä, M., Berndt, T., Petäjä, T., Brus, D., Vanhanen, J., Stratmann, F., Patokoski, J., Mauldin III, R. L.,
- 390 Hyvärinen, A.-P., and Lihavainen, H.: The role of sulfuric acid in atmospheric nucleation, Science,
- 391 327, 1243-1246, 2010.
- 392 Stone, D., Whalley, L. K., and Heard, D. E.: Tropospheric OH and HO 2 radicals: field
- measurements and model comparisons, Chemical Society Reviews, 41, 6348-6404, 2012.





- 394 Takegawa, N., Seto, T., Moteki, N., Koike, M., Oshima, N., Adachi, K., Kita, K., Takami, A., and
- 395 Kondo, Y.: Enhanced new particle formation above the marine boundary layer over the Yellow
- 396 Sea: Potential impacts on cloud condensation nuclei, Journal of Geophysical Research:
- 397 Atmospheres, 125, e2019JD031448, 2020.
- 398 Tham, Y. J., Sarnela, N., Iyer, S., Li, Q., Angot, H., Quéléver, L. L., Beck, I., Laurila, T., Beck, L. J., and
- 399 Boyer, M.: Widespread detection of chlorine oxyacids in the Arctic atmosphere, Nature
- 400 Communications, 14, 1769, 2023.
- 401 Thompson, C., Shepson, P., Liao, J., Huey, L., Apel, E., Cantrell, C., Flocke, F., Orlando, J., Fried, A.,
- 402 and Hall, S.: Interactions of bromine, chlorine, and iodine photochemistry during ozone
- depletions in Barrow, Alaska, Atmospheric Chemistry and Physics, 15, 9651-9679, 2015.
- 404 van Pinxteren, M., Fomba, K. W., van Pinxteren, D., Triesch, N., Hoffmann, E. H., Cree, C. H.,
- 405 Fitzsimons, M. F., von Tümpling, W., and Herrmann, H.: Aliphatic amines at the Cape Verde
- 406 Atmospheric Observatory: Abundance, origins and sea-air fluxes, Atmos. Environ., 203, 183-195,
- 407 2019.
- 408 Williamson, C. J., Kupc, A., Axisa, D., Bilsback, K. R., Bui, T., Campuzano-Jost, P., Dollner, M., Froyd,
- 409 K. D., Hodshire, A. L., and Jimenez, J. L.: A large source of cloud condensation nuclei from new
- particle formation in the tropics, Nature, 574, 399-403, 2019.
- Wood, R.: Stratocumulus clouds, Monthly weather review, 140, 2373-2423, 2012.
- 412 Xie, H.-B., Elm, J., Halonen, R., Myllys, N., Kurten, T., Kulmala, M., and Vehkamaki, H.: Atmospheric
- 413 fate of monoethanolamine: enhancing new particle formation of sulfuric acid as an important
- removal process, Environmental science & technology, 51, 8422-8431, 2017.
- 415 Yao, L., Garmash, O., Bianchi, F., Zheng, J., Yan, C., Kontkanen, J., Junninen, H., Mazon, S. B., Ehn,
- 416 M., and Paasonen, P.: Atmospheric new particle formation from sulfuric acid and amines in a
- 417 Chinese megacity, Science, 361, 278-281, 2018.
- 418 Yin, R., Yan, C., Cai, R., Li, X., Shen, J., Lu, Y., Schobesberger, S., Fu, Y., Deng, C., and Wang, L.: Acid
- 419 —base clusters during atmospheric new particle formation in urban Beijing, Environmental Science
- 420 & Technology, 55, 10994-11005, 2021.
- 421 Yu, H., Ren, L., Huang, X., Xie, M., He, J., and Xiao, H.: lodine speciation and size distribution in
- 422 ambient aerosols at a coastal new particle formation hotspot in China, Atmospheric Chemistry
- 423 and Physics, 19, 4025-4039, 2019.
- Zhang, R., Khalizov, A., Wang, L., Hu, M., and Xu, W.: Nucleation and growth of nanoparticles in
- 425 the atmosphere, Chemical reviews, 112, 1957-2011, 2012.
- Zhang, R., Suh, I., Zhao, J., Zhang, D., Fortner, E. C., Tie, X., Molina, L. T., and Molina, M. J.:
- 427 Atmospheric new particle formation enhanced by organic acids, Science, 304, 1487-1490, 2004.
- 428 Zheng, G., Wang, Y., Wood, R., Jensen, M. P., Kuang, C., McCoy, I. L., Matthews, A., Mei, F.,
- 429 Tomlinson, J. M., and Shilling, J. E.: New particle formation in the remote marine boundary layer,
- 430 Nature communications, 12, 527, 2021.
- 431 Zhu, Y., Li, K., Shen, Y., Gao, Y., Liu, X., Yu, Y., Gao, H., and Yao, X.: New particle formation in the
- 432 marine atmosphere during seven cruise campaigns, Atmospheric Chemistry and Physics, 19,
- 433 89-113, 2019.

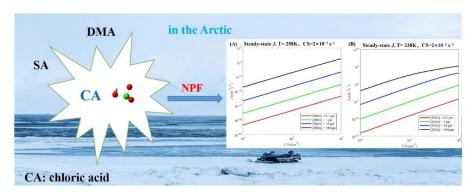
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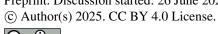
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Graphical Abstract







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466	Figure Captions
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468	Figure 1. Identified lowest free energy structures of the $(CA)_{1-4}(DMA)_{1-4}$ clusters at
469	the DLPNO-CCSD(T)/aug-cc-pVTZ// ω B97X-D/6-31++G(d,p) level of theory. The
470	red, blue, gray, green and white balls represent oxygen, nitrogen, carbon, chlorine and
471	hydrogen atoms, respectively. The dashed white and black lines indicate hydrogen
472	and halogen bonds, respectively.
473	Figure 2. The formation free energy (ΔG) (in kcal mol ⁻¹) of (A) (CA) ₁₋₄ (DMA) ₁₋₄ and
474	(B) $(CA)_{l-4}(SA)_{l-4}$ clusters at the
475	$DLPNO\text{-}CCSD(T)/aug\text{-}cc\text{-}pVTZ//\omega B97X\text{-}D/6\text{-}31\text{+++}G(d,p) level of theory. The supervision of the supervisi$
476	calculations are performed at 278 K and 1 atm.
477	Figure 3. Evaporation rates for (A) (CA) ₁₋₄ (DMA) ₁₋₄ and (B) (CA) ₁₋₄ (SA) ₁₋₄ clusters





- clusters at 278 K and 1 am.
- Figure 4. Simulated steady-state CA dimer concentration $\sum [(CA)_2]$ (cm⁻³) (A) and
- 480 the cluster formation rates J (cm⁻³ s⁻¹) out of the simulation systems (B) as a function
- 481 of [CA] at 278 K.
- Figure 5. The simulated cluster formation rate $J \text{ (cm}^{-3} \text{ s}^{-1})$ of the CA-DMA system at
- different temperatures (A) 258, and (B) 238 K; $[CA] = 10^6 10^8$ molec. cm⁻³; [DMA]
- 484 = 0.1, 1, 10, and 100 ppt; and $CS = 2 \times 10^{-3} \text{ s}^{-1}$.
- Figure 6. (A) Main clustering pathways of (CA)₁₋₄(DMA)₁₋₄ clusters at 278 K, [CA]
- 486 = 10^6 cm⁻³, and [DMA] = 1 ppt. (B) Main clustering pathways of $(CA)_{1-4}(SA)_{1-4}$
- clusters at 278 K, [CA] = 10^6 cm⁻³, and [SA] = 10^6 cm⁻³.





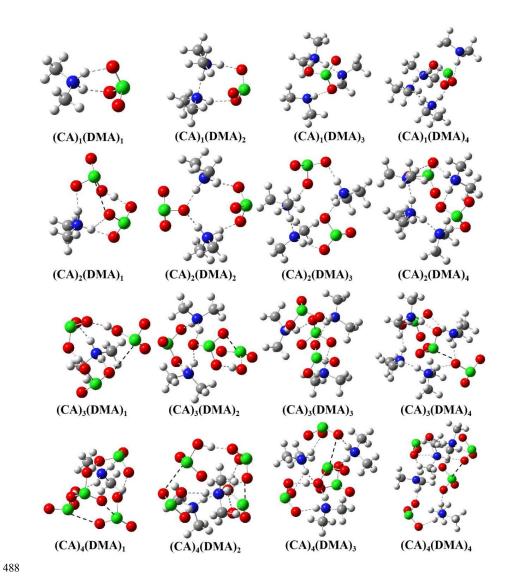
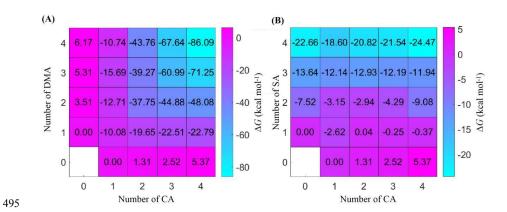


Figure 1. Identified lowest free energy structures of the $(CA)_{1-4}(DMA)_{1-4}$ clusters at the DLPNO-CCSD(T)/aug-cc-pVTZ// ω B97X-D/6-31++G(d,p) level of theory. The red, blue, gray, green and white balls represent oxygen, nitrogen, carbon, chlorine and hydrogen atoms, respectively. The dashed white and black lines indicate hydrogen and halogen bonds, respectively.







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Figure 2. The formation free energy (ΔG) (in kcal mol⁻¹) of (A) (CA)₁₋₄(DMA)₁₋₄ and

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at the

 $DLPNO\text{-}CCSD(T)/aug\text{-}cc\text{-}pVTZ//\omega B97X\text{-}D/6\text{-}31\text{+++}G(d,p) \quad level \quad of \quad theory. \quad The$

calculations are performed at 278 K and 1 atm.

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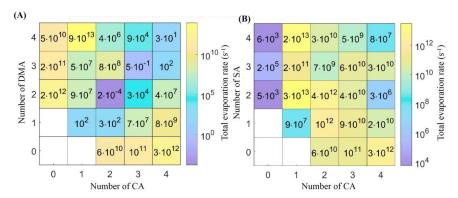


Figure 3. Evaporation rates for (A) $(CA)_{1-4}(DMA)_{1-4}$ and (B) $(CA)_{1-4}(SA)_{1-4}$ clusters

clusters at 278 K and 1 am.





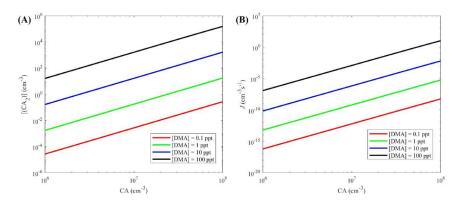


Figure 4. Simulated steady-state CA dimer concentration $\sum[(CA)_2]$ (cm⁻³) (A) and the cluster formation rates J (cm⁻³ s⁻¹) out of the simulation systems (B) as a function of [CA] at 278 K.





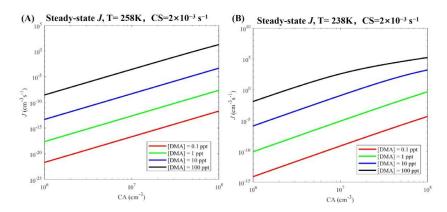
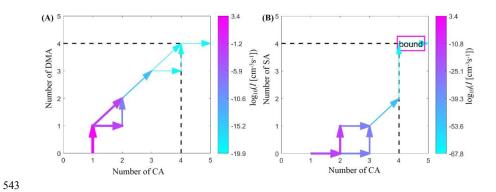


Figure 5. The simulated cluster formation rate J (cm⁻³ s⁻¹) of the CA-DMA system at different temperatures (A) 258, and (B) 238 K; [CA] = $10^6 - 10^8$ molec. cm⁻³; [DMA]

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Figure 6. (A) Main clustering pathways of (CA)₁₋₄(DMA)₁₋₄ clusters at 278 K, [CA]

546 = 10^6 cm⁻³, and [DMA] = 1 ppt. (B) Main clustering pathways of $(CA)_{1-4}(SA)_{1-4}$

547 clusters at 278 K, $[CA] = 10^6 \text{ cm}^{-3}$, and $[SA] = 10^6 \text{ cm}^{-3}$.

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