

Quantifying CO emissions from boreal wildfires by assimilating TROPOMI and TCCON observations

Sina Voshtani¹, Dylan B. A. Jones¹, Debra Wunch¹, Drew C. Pendergrass², Paul O. Wennberg^{3,4}, David F. Pollard⁵, Isamu Morino⁶, Hirofumi Ohyama⁶, Nicholas M. Deutscher⁷, Frank Hase⁸, Ralf Sussmann⁹, Damien Weidmann¹⁰, Rigel Kivi¹¹, Omaira García¹², Yao Té¹³, Jack Chen¹⁴, Kerry Anderson^{14,15}, Robin Stevens^{16,17,18}, Shobha Kondragunta¹⁹, Aihua Zhu²⁰, Douglas Worthy²¹, Senen Racki²², Kathryn McKain²³, Maria V. Makarova²⁴, Nicholas Jones⁷, Emmanuel Mahieu²⁵, Andrea Cadena-Caicedo²⁶, Paolo Cristofanelli²⁷, Casper Labuschagne²⁸, Elena Kozlova²⁹, Thomas Seitz³⁰, Martin Steinbacher³⁰, Reza Mahdi³¹, and Isao Murata³²

¹Department of Physics, University of Toronto, Toronto, Ontario, Canada

²School of Engineering and Applied Sciences, Harvard University, Cambridge, Massachusetts, USA

³Division of Geological and Planetary Sciences, California Institute of Technology, Pasadena, CA, USA

⁴Division of Engineering and Applied Science, California Institute of Technology, Pasadena, CA, USA

⁵New Zealand Institute for Earth Science Limited, Lauder, NZ

⁶National Institute for Environmental Studies (NIES), Onogawa 16-2, Tsukuba, Ibaraki, Japan

⁷Centre for Atmospheric Chemistry, School of Earth, Atmospheric and Life Sciences, University of Wollongong, Wollongong, Australia

⁸Institute of Meteorology and Climate Research (IMK-ASF), Karlsruhe Institute of Technology (KIT), Karlsruhe, Germany

⁹Karlsruhe Institute of Technology (KIT), IMK-IFU, Garmisch-Partenkirchen, Germany

¹⁰Space Science and Technology Department, STFC Rutherford Appleton Laboratory, Didcot, OX11 0QX, UK

¹¹Space and Earth Observation Centre, Finnish Meteorological Institute, Sodankylä, Finland

¹²Izaña Atmospheric Research Centre (IARC), State Meteorological Agency of Spain (AEMET), Santa Cruz de Tenerife, Spain

¹³Sorbonne Université, CNRS, MONARIS, UMR8233, F-75005 Paris, France

¹⁴Air Quality Research Division, Environment and Climate Change Canada, Toronto, Ontario, Canada

¹⁵Natural Resources Canada (emeritus)

¹⁶Department of Chemistry, Faculty of Arts and Sciences, Université de Montréal, Montréal, Quebec, Canada

¹⁷School of Public Health, Université de Montréal, Montréal, Quebec, Canada

¹⁸Climate Research Division, Environment and Climate Change Canada, Victoria, British Columbia, Canada

¹⁹Center for Satellite Application and Research, NOAA/NESDIS, College Park, USA

²⁰IM Systems Group Inc., College Park, USA

²¹Climate Research Division, Environment and Climate Change Canada, Toronto, Ontario, Canada

²²Climate Chemistry Measurements and Research, Environment and Climate Change Canada, Toronto, Ontario, Canada

²³NOAA Global Monitoring Laboratory, Boulder, CO, USA

²⁴Department of Atmospheric Physics, Faculty of Physics, St. Petersburg State University, Saint Petersburg, Russia

²⁵Institute of Astrophysics and Geophysics, University of Liège, Liège 4000, Belgium

²⁶Instituto de Ciencias de la Atmósfera y Cambio Climático Universidad Nacional Autónoma de México, UNAM, Mexico City, Mexico

²⁷National Research Council of Italy, Institute of Atmospheric Sciences and Climate (ISAC), Bologna, Italy

²⁸South African Weather Service, Stellenbosch, South Africa

²⁹University of Exeter, Hatherly Labs, Exeter, EX4 4PS, UK

³⁰Empa, Swiss Federal Laboratories for Materials Science and Technology, CH-8600 Dübendorf, Switzerland

³¹GAW Bukit Kototabang Station, BMKG, Indonesia

³²Graduate School of Environmental Studies, Tohoku University, Sendai, Japan

Correspondence to: Sina Voshtani (voshtani.sina@gmail.com; sina.voshtani@utoronto.ca)

Abstract. We perform a global inverse modelling analysis to quantify biomass burning emissions of carbon monoxide (CO) from the extreme wildfires in Canada between May and September 2023. Using the GEOS-Chem model, we assimilated

observations at 3-day temporal and $2^\circ \times 2.5^\circ$ horizontal resolution from the Tropospheric Monitoring Instrument (TROPOMI) separately and then jointly with Total Carbon Column Observing Network (TCCON) measurements. We also evaluated prior emissions from the Quick Fire Emissions Dataset (QFED), Blended Global Biomass Burning Emissions Product eXtended (GBBEPx), Global Fire Assimilation System (GFAS), and Canadian Forest Fire Emissions Prediction System (CFFEPS). The assimilation of TROPOMI-only measurements estimated posterior North America emissions for QFED, GBBEPx, GFAS, and CFFEPS of 110.4 ± 20 , 112.8 ± 20 , 127.2 ± 17 , and 125.6 ± 18 Tg CO compared to prior estimates of 37.1, 42.7, 91.0, and 90.2 Tg CO, respectively. The joint assimilation of TROPOMI+TCCON reduced the posterior 1σ uncertainty on the North American emission estimates by up to about 30%, while showing only a modest impact ($< 5\%$) on the mean estimate of the inferred emissions. An evaluation against independent measurements reveals that adding TCCON data increases the correlations and slightly lowers the biases and standard deviations. Additionally, including an experimental TCCON product at East Trout Lake with higher surface sensitivity, we find better agreement of the assimilation results with nearby in situ tall tower and aircraft measurements. This highlights the potential importance of vertical sensitivity in these experimental data for constraining local surface emissions. Our results demonstrate the complementarity of the greater temporal coverage provided by TCCON with the spatial coverage of TROPOMI when these data are jointly assimilated.

1 Introduction

Biomass burning (BB) from wildfires is a major source of carbon emissions released into the atmosphere with large climate and air quality impacts, exerting a significant influence on human health, ecosystems, and the environment (Cascio, 2018; Chen et al., 2017; O'Neill et al., 2021; Wu et al., 2022). Over the past few years, wildfires have become more frequent and destructive in different regions of the world (Jegasothy et al., 2023; Mataveli et al., 2024; You and Xu, 2023). More specifically, in Canada in 2023, the total area burned by wildfires surpassed the previous record in 1989 ($75,596 \text{ km}^2$) by nearly a factor of 2.5 while the amount of emitted carbon also dramatically increased by more than 11 times compared to the 1998–2022 average (Jain et al., 2024; Jones et al., 2024; Kolden et al., 2024). Therefore, reactive trace gases (e.g., CO) and greenhouse gases (GHGs) (e.g., CH_4 and CO_2) had an unprecedented increase during the 2023 wildfire emissions (Byrne et al., 2024). In addition to perturbing the carbon budget, these emissions also have implications for air quality. However, obtaining reliable estimates of wildfire emissions is a challenging task due to several factors, such as the episodic and localized nature of these emissions (Sokolik et al., 2019; Zhao et al., 2025). Here, we focus on estimating emissions from carbon monoxide (CO) from wildfires during the summer 2023. With a lifetime of up to several months (Holloway et al., 2000), which is sufficiently long to track long-range transport on intercontinental scales, CO is an ideal tracer of combustion. CO plays an important role in both air quality and climate as it is a precursor of ozone (O_3) and the dominant sink of the hydroxyl radical (OH), which is the main atmospheric oxidant (Aschi and Largo, 2003; Fowler et al., 2008). Observations of CO have provided information on combustion sources on a range of scales, from urban to regional and global scales (Borsdorff et al., 2020; Cristofanelli et al., 2024; Pommier et al., 2013; Schneising et al., 2020; Tang et al., 2019).

CO emissions from wildfires can be estimated either from bottom-up or top-down approaches. In the bottom-up approach, emissions are represented as the product of an emission factor, which is the amount of trace gas emitted per unit of fuel consumed, and the amount of dry matter burned. Bottom-up inventories use either observations of burned area to determine the mass of dry matter burned (Liu et al., 2024; van der Werf et al., 2017; Wiedinmyer et al., 2023) or estimates of fire radiative power (FRP) to quantify the rate of fuel consumption (Filizzola et al., 2023; Kaiser et al., 2012). There are typically large uncertainties in these inventories (Andreae, 2019; Hundal et al., 2024) arising from discrepancies in the emission factors and the estimated mass of dry matter burned, resulting in significant differences in emission estimates (Chen et al., 2022; Nguyen et al., 2023; Saikawa et al., 2017; Zhang et al., 2023). The top-down approach makes use of CO observations to optimize emissions through an inverse modelling method, but this approach depends on the use of an atmospheric chemistry-transport model and a priori emission estimates, which are typically obtained from a bottom-up inventory.

Over the past two decades, global CO observations have been provided by several satellite sensors, including the Measurements of Pollution in the Troposphere (MOPITT) instrument (Deeter et al., 2003; Edwards et al., 2006), the Infrared Atmospheric Sounding Interferometer (IASI) (Pope et al., 2021; Turquety et al., 2004), and the Tropospheric Emission Spectrometer (TES) (Lopez et al., 2008), and these data have been used in numerous inverse modelling studies to quantify CO emissions (e.g., Kasibhatla et al., 2002; Arellano Jr. et al., 2006; Warner et al., 2007; Jones et al., 2009; Kopacz et al., 2010; Miyazaki et al., 2012; Miyazaki, Eskes, and Sudo 2015; Jiang et al., 2017; Zheng et al., 2019). However, large discrepancies between the inversion results have been reported, which may arise from differences between spatiotemporal coverage of the observations, the vertical sensitivity of the measurements, and observation biases (Deeter et al., 2015; Jiang et al., 2017; Jones et al., 2009; Miyazaki et al., 2015; Warner et al., 2010). Nonetheless, a few recent studies attempted to address some of the challenges by reducing potential biases in the model (Gaubert et al., 2023; Miyazaki et al., 2020) or by improving the quality of the assimilated data (Tang et al., 2024). The Tropospheric Monitoring Instrument (TROPOMI) (Borsdorff et al., 2018), launched in 2017, has provided CO retrievals with improved accuracy, higher spatial resolution, and significantly greater observational coverage (Landgraf et al., 2016; Schneising et al., 2020). In particular, it offers higher sensitivity to near-surface CO compared to earlier thermal infrared measurements from instruments such as IASI and TES. These factors make it well-suited for inverse modelling of CO emissions, as demonstrated in many recent studies (Borsdorff et al., 2023; Byrne et al., 2024; Goudar et al., 2023; Griffin et al., 2024; Inness et al., 2022; Shahrokhi et al., 2023; Stockwell et al., 2022; Wan et al., 2023).

Measurements from surface in situ networks and aircraft campaigns have been used for CO trend determination (Patel et al., 2024) and CO inversion studies in the past (Palmer et al., 2003; Yumimoto & Uno, 2006; Koohkan & Bocquet, 2012; Tang et al., 2013; Feng et al., 2020). However, mainly due to limited spatiotemporal coverage and/or vertical distribution, they are typically incapable of sufficiently constraining emission estimates on fine spatial scales, and therefore, model errors, such as those from vertical transport, the OH field, and the a priori emissions, can significantly impact the inferred emission estimates (Hooghiemstra et al., 2011). Only a few studies have attempted to use both satellite and surface observations together to exploit the complementarity of these observations to reduce the influence of errors, such as those that arise from the

sensitivity to the vertical distribution of CO (Tang et al., 2022) and long-range transport (Kim et al., 2024). However, most of these studies focused on inversions over a limited area, where sufficient surface CO observations are available. There are also ground-based total column measurements from networks such as the Total Carbon Column Observing Network (TCCON, Wunch et al. 2011), which was designed in part to validate satellite observations (Borsdorff et al., 2019; Bukosa et al., 2023; Hedelius et al., 2021; Sha et al., 2021; Tang et al., 2024). The TCCON provides time-resolved and accurate column-averaged dry-air mole fractions of CO (XCO) under sunny skies. Although TCCON observations are spatially sparse, they are of high temporal density, and therefore could provide valuable information in constraining episodic CO emissions from wildfires. However, a standard method of integrating TCCON measurements with satellite data in a data assimilation or inversion system is still lacking, as most current studies assimilate satellite data, while reserving the TCCON data to evaluate the performance of the assimilation.

In this study, we quantify biomass burning CO emissions between May and September 2023 using the CHEMistry and Emissions REanalysis Interface with Observations (CHEEREIO) assimilation toolkit (Pendergrass et al., 2023), which employs the GEOS-Chem model and an ensemble Kalman filter (EnKF) scheme. We conduct a global analysis, but our focus is on quantifying boreal emissions associated with the 2023 fires in Canada. We jointly assimilate TCCON and TROPOMI data and conduct a comparison with a TROPOMI-only assimilation to assess the added value of TCCON observations in the assimilation and to determine the additional constraints that TCCON data provide for optimizing CO emissions from localized and episodic wildfires. We include two distinct types of TCCON data with different vertical sensitivities in our inversion, while using independent total column and in situ surface and aircraft vertical profile observations to characterize the success of our analysis. Additionally, we evaluate the following three global and one regional biomass burning inventories in the context of the assimilation: the Quick Fire Emissions Dataset (QFED) (Koster et al., 2015), the Blended Global Biomass Burning Emissions Product eXtended (GBBEPx) (Zhang et al. 2012; 2019), the Global Fire Assimilation System (GFAS) product (Kaiser et al. 2012; Di Giuseppe et al. 2021), and the Canadian Forest Fire Emissions Prediction System (CFFEPS) (Chen et al., 2019).

We begin in Section 2 with a description of the observations and model configurations, including the a priori emissions used in the inversion. Section 3 presents the inversion methods, the main assumptions, and sensitivity experiments using simulated observations to tune the inversion performance. Section 4 provides the main results and discussion, and finally, the study concludes with a few summary points and a suggestion for future works.

2 Observations and model

This section describes the observational datasets and modelling framework and inputs used in this study. We first describe in Sections 2.1-2.2 the two types of observations that are assimilated in the inversion framework: (i) TROPOMI satellite CO products and (ii) TCCON ground-based CO measurements. Then, in Section 2.3, we present several independent datasets (not assimilated) used for evaluation, including the Network for the Detection of Atmospheric Composition (NDACC,

De Mazière et al., 2018) ground-based total column observations, in situ surface CO measurements from the World Data Centre for Greenhouse Gases (WDCGG) network and from Environment and Climate Change Canada’s (ECCC) tall tower at East Trout Lake (ETL), and vertical profiles from in situ aircraft measurements. The ~~same~~ assimilated TCCON data are also used for comparisons between different experiments, although they are no longer independent information for the joint inversion. Finally, Section 2.4 provides a description of the GEOS-Chem model and emissions inventories that are used as the a priori estimate in our inversion ~~2-s~~ setup. The priors include three global inventories, including QFED, GBBEPx, and GFAS, and one regional inventory, CFFEPS, for North America.

2.1 TROPOMI

The TROPOMI instrument is on board the Copernicus Sentinel 5 Precursor (S5P) satellite, which was launched in October 2017. Total column abundances of carbon monoxide (XCO) are retrieved from spectra measured in the shortwave infrared (SWIR) band at 2305–2385 nm, with daily global coverage, a local overpass solar times of 13:30 UTC, and high spatial resolution of $5.5 \times 7 \text{ km}^2$ (Veefkind et al., 2012). We use the operational XCO product publicly available from the European Space Agency (ESA) Sentinel-5P data hub at <https://scihub.copernicus.eu/> (last access: 4 June 2024) (Landgraf, 2019), with a reported bias of better than 15 ppb in comparison with TCCON GGG2014 data product (Sha et al., 2021). The XCO data are published together with the total column averaging kernels to account for the sensitivity of the retrieved total column to the true atmosphere, thus, they can be used along with a priori vertical profiles to obtain model-equivalent total CO columns representing the observed data (Apituley et al., 2018). The TROPOMI retrieval algorithm provides clear-sky and cloudy observations over land and ocean (Borsdorff et al., 2019), however, we only use measurements with a quality flag equal to and greater than 0.7 to ensure high-quality data obtained under cloud-free or low cloud conditions. As shown in the Fig. 1a, we exclude TROPOMI observations poleward of 60° , primarily to avoid biases due to low surface albedo in the SWIR from snow cover (Hasekamp et al., 2022; Lorente et al., 2021) and biases due to the stratosphere in the chemical transport model (CTM) affecting the inversion performance (Turner et al., 2015). Since the horizontal resolution of the TROPOMI data is substantially higher than the GEOS-Chem model resolution used in this study ($2^\circ \times 2.5^\circ$), the observations are not spatially representative for the model grid cells, resulting in a large representativeness error ~~within~~ the assimilation process. To overcome this, we aggregate the observations into so-called super-observations before using them in the assimilation. In fact, for the duration of the study between May and September 2023, we compute the error-weighted median average of measurements within each grid cell, where each measurement is weighted by the inverse of its reported error standard deviation (Eskes et al., 2003; Miyazaki et al., 2012). To account for error reduction because of averaging, we follow a similar method as Pendergrass et al. (2023) to compute the associated super-observation errors. This includes average of individual measurement errors and assumptions on error correlations and transport errors. The relative weight of the super-observation error to the prior error is then estimated through parameter tuning in the inversion system to ensure robustness to possible error misspecification (see Section 3 and Appendices A and B). Finally, a total of 1,744,682 number of observations are processed.

Several previous studies evaluated TROPOMI XCO observations and found reasonable agreement with satellite and ground-based measurements. For example, Sha et al. (2021)⁷ reported a bias of $2.45 \pm 3.38\%$ against the unscaled TCCON and a bias of $6.50 \pm 3.45\%$ against NDACC, which remains within the range of TROPOMI's precision and accuracy. In addition, the TROPOMI validation reports (<https://mpc-vdaf.tropomi.eu/>, last access: 28 June 2024)³⁵ Lambert et al., 2024) shows that operational TROPOMI XCO data are in good agreement with collocated measurements from NDACC, TCCON, and the Collaborative Carbon Column Observing Network (COCCON)³⁵ Alberti et al., 2022; Frey et al., 2019) monitoring networks.

2.2 TCCON

TCCON (<https://www.tccon.caltech.edu/>) is a ground-based network of solar-viewing Fourier transform spectrometers (FTS) that collect atmospheric transmission spectra every 2–3 minutes. The spectra range covers the near and short-wave infrared region, and measurements collected under clear-sky conditions are used to retrieve column-averaged dry-air mole fractions of trace gases, including carbon monoxide (i.e., XCO) (Wunch et al., 2011). We use data from 15 sites around the world, shown in Fig. 1 and listed in Table 1, derived from the standard GGG2020 retrieval software (Laughner et al., 2024). The difference between the GGG2020 and GGG2014 XCO retrieval data is 6.3 ppb, with GGG2020 larger than GGG2014. This is because TCCON XCO data are no longer scaled to the WMO trace gas scale (Wunch et al., 2025). The accuracy and precision of the standard TCCON XCO product is reported to be around 8 ppb. These data are publicly available and can be accessed via <https://tccondata.org/> (last access: 1 June 2024). In addition to the standard XCO retrievals, which use spectra measured using an InGaAs detector and a CO window centred at 4290 cm^{-1} , we use retrievals of XCO available from spectra collected at the East Trout Lake TCCON station from an additional InSb detector. The spectral range of the InSb detector includes two mid-infrared windows (centred at 2111 cm^{-1} and 2160 cm^{-1}) that contain strong CO absorption features that result in an XCO retrieval with markedly different averaging kernels (orange) from the standard XCO retrieval (blue), as shown in Fig. 1c-d. These mid-infrared spectral windows were used in a previous study together with the standard TCCON CO window to extract vertical information from the TCCON measurements (Parker et al., 2023). The XCO retrievals from the InSb spectra have higher sensitivity to the surface and lower sensitivity to the higher altitudes than the standard XCO retrievals. The impact of including these mid-infrared XCO retrievals on the inversion performance to constrain CO emissions is discussed in Section 4.2.2.

We filter all TCCON datasets to include only data with a quality flag = 0 (i.e., the highest quality data). To prepare for the assimilation, first, all the measurements are aggregated in time, based on the model output hourly timestep, weighted by the measurement reported errors. This produced 213,784 quality-controlled data points that were then mapped on the GEOS-Chem grid resolution, providing 19,733 median hourly averaged observations for the period of May–September 2023 in this study.

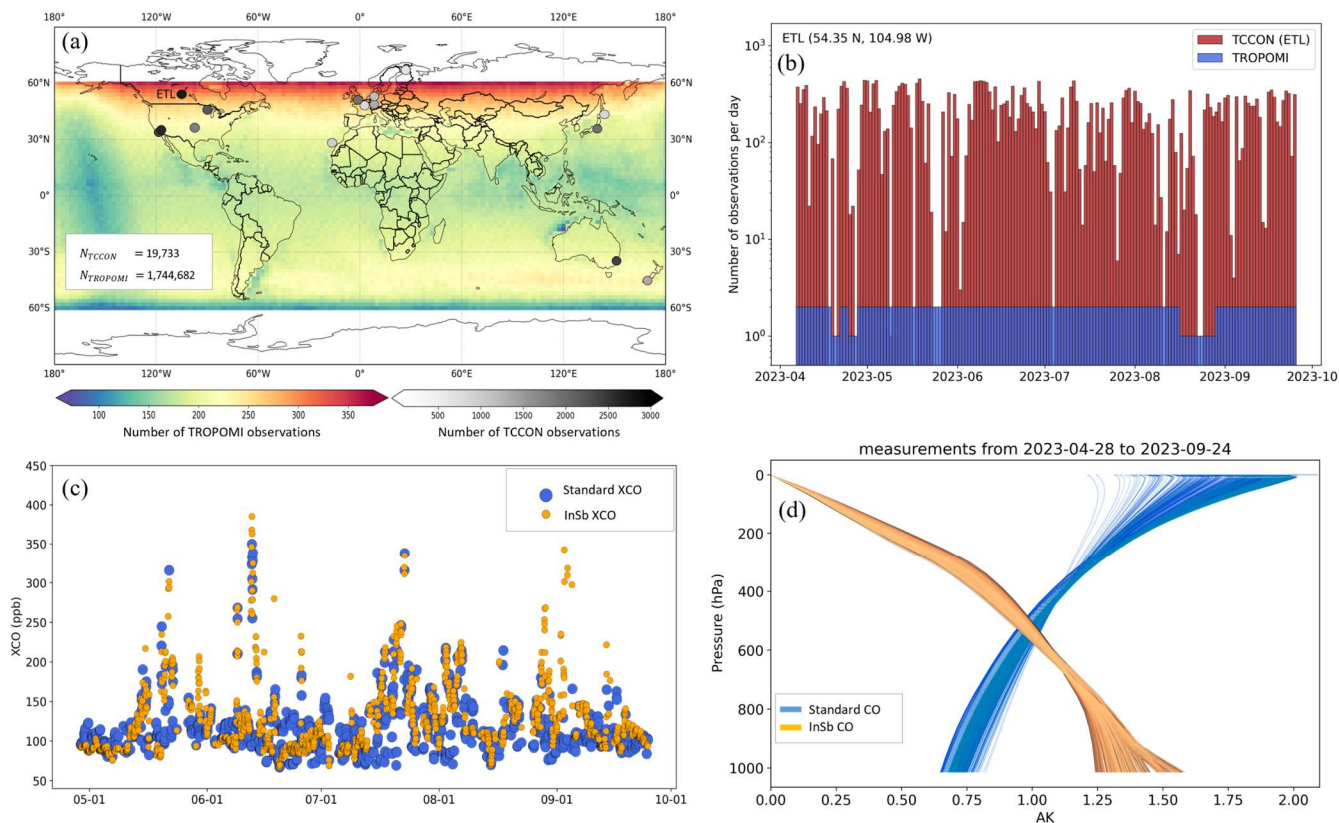


Figure 1: (a) The number of quality-controlled and aggregated TROPOMI and TCCON observations (XCO) used in the global inversion as described in Sections 2.1 and 2.2. (b) The variability of the number of non-aggregated TROPOMI (blue) and TCCON (red) observations at the East Trout Lake (ETL) TCCON station from April-September 2023. (c) Time series (MM-DD) of XCO column retrievals from standard TCCON GGG2020 data (blue) and XCO measurements from the InSb detector (orange) at East Trout Lake (ETL) between May-September 2023. (d) Column averaging kernels for the standard TCCON GGG2020 XCO (blue) and column averaging kernels for XCO measurements derived from an alternative CO absorption window on the ETL InSb detector (orange) at East Trout Lake (ETL) between May-September 2023.

2.3 NDACC and in situ surface and aircraft data

This study employs two independent ground-based data sources for validation purpose. We utilize measurements from NDACC, a global network of ground-based stations equipped with Fourier transform infrared (FTIR) spectrometers that provide long-term total column measurements of XCO (De Mazière et al., 2018). NDACC XCO measurements, similar to those from TCCON (i.e., same spectral domain as the InSb TCCON data), are of high-quality and well-suited for validating models, satellite observations, and assimilation system performance (Kerzenmacher et al., 2012; Lutsch et al., 2020; Sha et al., 2021). In this study, we include mid-infrared NDACC total column data from seven stations covering the study period (see Table 1). The data are publicly available at <http://www.ndacc.org> (last access: 30 July 2024).

Additionally, continuous and discrete surface in situ CO measurements obtained from WDCGG serve as a second independent dataset to evaluate surface CO concentrations obtained from our experiments (see Table 2). In situ measurements

225 compiled by the WDCGG have been widely used in previous inverse modelling studies for validating results and testing model
performance (Chevallier et al., 2011; Jiang et al., 2017; Miyazaki et al., 2020; Tanimoto et al., 2008). We use the archived data
from nine sites over the study period, which are publicly available and accessed from <https://gaw.kishou.go.jp/> (last access: 1
July 2024). We also use in situ tall tower measurements from the ETL site provided by ECCO (Chen et al., 2014) to assess the
impact of using the XCO retrievals from the InSb spectra on the inversion results. The evaluation results are presented in
230 Section 4.2.2.

We use in situ aircraft CO measurements from the National Oceanic and Atmospheric Administration (NOAA) air
sampling network (McKain et al., 2024) taken as another independent source to evaluate our inversion results. The data product
is freely available to public via <https://gml.noaa.gov/ccgg/aircraft/>. The aircraft program aims to capture temporal variability
(i.e., seasonal and interannual changes) of the greenhouse gases in the lower atmosphere. Dry air mole fractions of CO are
235 measured using flask air samples at different fixed altitude levels. It provides measurements at different sites across the United
States and Canada and at different altitudes, descending from a maximum of 8000 m to the lowest sampling level at ~750 m
(a.s.l). These data have been commonly used in previous studies to explore the large-scale changes in horizontal and vertical
distribution of CO and greenhouse gases (Sweeney et al., 2015), to serve as benchmark for validating forward and inverse
modelling analysis (Stephens et al., 2007; Yang et al., 2007), and to calibrate remote sensing retrievals (Wunch et al., 2010).
240 Focusing on the impact of the experimental TCCON InSb data used in the inversion to constrain surface CO emissions, we
use aircraft profiles at ETL during multiple time events (details are discussed in Section 4.2.2). Table 1 shows the list and
geographical information of all observations used for evaluation.

Table 1: List of ground network, surface in situ, aircraft, and tall tower measurements used in this study between May and September 2023.

Measurement	Site (ID)	Latitude	Longitude	Altitude (km a.s.l)	Reference
TCCON	Sodankylä, Finland	67.4° N	26.6° E	0.188	(Kivi et al., 2022)
TCCON	Karlsruhe, Germany	49.1° N	8.4° E	0.116	(Hase et al., 2023)
TCCON	Garmisch, Germany	47.4° N	11.1° E	0.740	(Sussmann and Rettinger, 2023)
TCCON	Paris, France	48.8° N	2.4° E	0.060	(Té et al., 2022)
TCCON	Harwell, UK	51.6° N	1.3° W	0.142	(Weidmann et al., 2023)
TCCON	Izana, Tenerife, Spain	28.3° N	16.5° W	2.370	(García et al., 2022)
TCCON	East Trout Lake, Canada	54.3° N	104.9° W	0.502	(Wunch et al., 2022)
TCCON	Lamont, USA	36.6° N	97.5° W	0.320	(Wennberg et al., 2022b)
TCCON	Park Falls, USA	45.9° N	90.3° W	0.440	(Wennberg et al., 2022c)
TCCON	Caltech, USA	34.1° N	118.1° W	0.230	(Wennberg et al., 2022a)
TCCON	Edwards, USA	34.0° N	117.8° W	0.700	(Iraci et al., 2022)
TCCON	Rikubetsu, Japan	43.5° N	143.8° E	0.380	(Morino et al., 2022a)
TCCON	Tsukuba, Japan	36.1° N	140.1° E	0.030	(Morino et al., 2022b)
TCCON	Wollongong, Australia	34.4° S	150.9° E	0.300	(Deutscher et al., 2023)
TCCON	Lauder, New Zealand	45.0° S	169.7° E	0.370	(Pollard et al., 2022)
NDACC	Tsukuba, Japan	36.1° N	140.1° E	0.030	(Morino et al., 2022b)
NDACC	Wollongong, Australia	34.4° S	150.9° E	0.300	(Jones et al., 2009)
NDACC	Lauder, New Zealand	45.0° S	169.7° E	0.370	(Bègue et al., 2024)
NDACC	Arrival Heights, Antarctica	77.8° S	66.67° E	0.184	(Smale et al., 2021)
NDACC	St. Petersburg, Russia	59.9° N	29.8° E	0.020	(Makarova et al., 2024)
NDACC	Jungfraujoch, Switzerland	46.5° N	7.9° E	3.580	(Zander et al., 2008)
NDACC	Altzomoni, Mexico	19.1° N	98.6° W	3.985	(Grutter et al., 2008)
In Situ	Bukit Kototabang (BKT), Indonesia	0.2° S	100.3° E	0.864	(Eko Cahyono et al., 2022), Reza Mahdi, BMKG
In Situ	Minamitorishima (MNM), Japan	24° N	153.9° E	0.007	(Takatsuji, S., 2024a)
In Situ	Ryori (RYO), Japan	39.3° N	141.8° E	0.260	(Takatsuji, S., 2024b)
In Situ	Yonagunijima (YON), Japan	24.4° N	123° E	0.030	(Takatsuji, S., 2024c)
In Situ	Capo Granitola (CGR), Italy	37.6° N	126° E	0.005	(Cristofanelli et al., 2017)
In Situ	Cape Point (CPT), South Africa	34.3° S	18.5° E	0.230	(Labuschagne et al., 2018)
In Situ	Cape Verde Atmospheric Observatory (CVO), Cabo Verde	16.9° N	24.9° W	0.010	(Kozlova, E. et al., 2021)
In Situ	Jungfraujoch (JFJ), Switzerland	46.5° N	7.9° E	3.580	(Hueglin et al., 2024), Martin Steinbacher, Empa
In Situ	Mt. Kenya (MKN), Kenya	0.06° S	37.3° E	3.678	(Kirago et al., 2023), Martin Steinbacher, Empa
Tall tower	East Trout Lake, Canada	54.3° N	104.9° W	0.502 ^a	(Chen et al., 2014), Douglas Worthy, ECCC
Aircraft	East Trout Lake, Canada	54.3° N	104.9° W	-	(McKain et al., 2024)

^a This is the surface altitude, and the measurement intake are at four levels (95, 55, 33, 22 m) installed on a 105 m SaskTel communication tower.

245 2.4 GEOS-Chem and prior estimates

The GEOS-Chem model (<http://www.geos-chem.org>, last access: 1 July 2024) is a global 3D CTM that uses assimilated meteorological observations as input from the NASA Global Modelling and Assimilation Office (GMAO). We use version 14.1.1 of the GEOS-Chem CTM driven by meteorological input from the Modern-Era Retrospective analysis for Research and

Applications, Version 2 (MERRA-2; Gelaro et al., 2017). The meteorological fields have a native resolution of $0.25^\circ \times 0.3125^\circ$ with 72 vertical levels from the surface to 0.01 hPa, which is degraded to $2^\circ \times 2.5^\circ$ horizontal grid and 47 vertical levels (Bey et al., 2001). For the purpose of global CO assimilation in this study, the linear CO-only simulation of GEOS-Chem, also known as “tagged CO”, is used with prescribed monthly mean OH fields from a 10-year archived full chemistry simulation based on version 14 of the model. The tagged CO simulation reduces the computation cost relative to the full-chemistry and has been widely applied in different studies in the past (Heald et al., 2004; Jiang et al., 2017; Jones et al., 2009; Kopacz et al., 2010; Lutsch et al., 2020; Tang et al., 2023; Wunch et al., 2019). Version 14.1.1 of the tagged CO simulation incorporated the improved secondary CO production scheme for the tagged CO simulation, which reduces the differences between full chemistry and tagged CO simulations, especially in regions strongly influenced by biogenic emissions and chemistry (Fisher et al., 2017). The biogenic source of CO in the full chemistry simulation is based on the oxidation of volatile organic compounds (VOCs) produced by the Model of Emissions of Gases and Aerosols from Nature (MEGAN version 2.1) inventory (Guenther et al., 2012).

For the results presented here, we specify fossil fuel emissions of CO from the Community Emissions Data System (CEDS) inventory (Hoesly et al., 2018). Biomass burning (BB) emissions are based on the four different BB inventories described below. These BB emissions have been used in various studies (Griffin et al., 2020; Jin et al., 2024; Li et al., 2020; Zhang et al., 2022), and are used as our prior emissions in the inversion analyses conducted here.

2.4.1 QFED

The Quick Fire Emissions Dataset version 2.5r1 (QFED v2.5r1) (Koster et al., 2015), is a global product of biomass burning emissions which was developed for the NASA GEOS model. It applies the fire radiative power (FRP) method with a cloud correction technique (Koster et al., 2015), where the location of fires and FRP are derived from the polar orbiting Moderate Resolution Imaging Spectroradiometer (MODIS) instrument onboard the NASA’s Terra and Aqua satellites. QFED provides daily emissions with a horizontal spatial resolution of $0.1^\circ \times 0.1^\circ$, and GEOS-Chem applies a climatological profile based on WRAP (Western Regional Air Partnership) method (WRAP, 2005) to distribute the emissions over the diurnal cycle. In the QFED implementation in GEOS-Chem, the setup of the plume injection height follows Fischer et al. (2014) and Travis et al. (2016), where the 65% of the biomass burning emissions are allocated to the planetary boundary layer (PBL) and the remaining 35% belongs to the free troposphere. The QFED data can be accessed from <http://geoschemdata.wustl.edu/ExtData/HEMCO/QFED/v2023-05/> (last access: 1 July 2024).

2.4.2 GBBEPx

The Blended Global Biomass Burning Emissions Product eXtended (GBBEPx v4, Zhang et al., 2012; 2019), developed by NOAA National Environmental Satellite, Data, and Information Service (NESDIS), produces daily global biomass burning emissions. GBBEPx blends information from QFED and fire emissions estimated from the Visible Infrared Imaging Radiometer Suite (VIIRS) instrument onboard Suomi National Polar-orbiting Partnership (SNPP) and Joint Polar Satellite

System (JPSS). VIIRS fire emissions are obtained using FRP derived from VIIRS data in an approach that is similar to the use of MODIS FRP data in QFED (Csiszar et al., 2016), but with a different fire detection scheme (Zhang et al., 2019). The blended GBBEPx emissions are produced daily at a resolution of $0.25^\circ \times 0.3125^\circ$, and the same profile is applied to distribute the emissions over the diurnal cycle as is used for QFED. The implementation of GBBEPx in GEOS-Chem assumes the same plume injection height scheme as for QFED. The GBBEPx v4 data can be accessed from <https://www.ospo.noaa.gov/pub/Blended/GBBEPx/> (last access: 1 July 2024).

2.4.3 GFAS

The Global Fire Assimilation System (GFAS v1.2, Kaiser et al. 2012; Di Giuseppe et al., 2021), utilized by the Copernicus Atmosphere Monitoring Service (CAMS), provides daily estimate of biomass burning emissions by assimilating FRP observations from MODIS instruments on the Terra and Aqua satellites. GFAS estimates emissions by conversion of FRP to the dry matter burned and the use of biome-specific emission factors. GFAS utilizes the vegetation-type prescribed by the Global Fire Emissions Database (GFED). The daily data are globally available at a resolution of $0.1^\circ \times 0.1^\circ$ from 2003 to the present time, which can be accessed from <https://ads.atmosphere.copernicus.eu/datasets/cams-global-fire-emissions-gfas?tab=overview>, (last access: 11 June 2024). In the GEOS-Chem simulations, we use the same diurnal cycle as used in QFED and GBBEPx. Additionally, GFAS provides information about the daily injection height (i.e., mean altitude of maximum injection (MAMI)) of the emissions. In GEOS-Chem, it is assumed that the emissions are injected uniformly from the surface to the MAMI.

2.4.4 CFFEPS

The Canadian Forest Fire Emissions Prediction System (CFFEPS v4.0, Chen et al., 2019) produces biomass burning emissions for North America using information from the Canadian Forest Service (CFS), Canadian Wildland Fire Information System (CWFIS), and meteorological inputs from ECCO's Global Environmental Multiscale (GEM) model. The product provides hourly fire emissions and smoke plume injection height at individual hotspot locations. In implementing the CFFEPS emissions in GEOS-Chem, the emissions were aggregated to the GEOS-Chem grid resolution with a weighted average plume height based on the CO₂ emission level. The CO₂ emission level was used in determining the injection height for all species in CFFEPS, even though here we focus only on the CO emissions. The plume injection height estimates from the FireWork plume rise model (Anderson et al., 2011; Chen et al., 2019) on which CFFEPS is based, have been validated against satellite-driven (e.g., TROPOMI) driven aerosol plume heights (Griffin et al., 2020).

3 Inversion methodology

The inversion framework utilizes the CHEEREIO assimilation toolkit (Pendergrass et al., 2023), which employs a localized ensemble transform Kalman filter (LETKF). A detailed description of the LETKF algorithm used is provided in Hunt

et al. (2007). We use CHEEREIO to derive optimized estimates of globally gridded emissions of CO between May and September 2023 at a spatial resolution of $2^\circ \times 2.5^\circ$ and a temporal resolution of three days by assimilation of TROPOMI satellite and TCCON observations. The solution state vector of emissions, $\mathbf{x}^a$, is given by

$$\overline{\mathbf{x}^a} = \overline{\mathbf{x}^b} + \gamma \mathbf{X}^b \tilde{\mathbf{P}}^a (\mathbf{Y}^b)^T \mathbf{R}^{-1} (\mathbf{y}^o - \overline{H(\mathbf{x}^b)}), \quad (1)$$

where the overbar represents the ensemble mean, $\mathbf{x}^b$ is the background state vector, $\mathbf{y}^o$ is the observation vector, $H(\mathbf{x}^b)$ is the model simulation of the observations with observation operator H , $\mathbf{R}$ is the observation error covariance matrix, $\tilde{\mathbf{P}}^a$ is the analysis error covariance matrix where tilde represents the ensemble space, $\mathbf{X}^b$ is the background perturbation matrix, $\mathbf{Y}^b$ is the observation perturbation matrix (see the detailed description of LETKF variables in Hunt et al., 2007). γ is a regularization factor used to prevent overfitting or underfitting to observations by balancing the relative influence of the a priori estimate and the measurements in the inversion. It serves as a pragmatic correction for uncertainties that are hard to quantify, such as unaccounted observation error correlations (Hakami et al., 2005; Lu et al., 2022; Voshtani et al., 2023).

To simulate the observations, we have developed an observation operator for each measurement type. The observation operator maps the model emission fields (i.e., state space) into the observation space as follow:

$$H(\mathbf{x}^b) = h_v [\mathbf{c}^{a\,priori} + \mathbf{A}(\mathbf{M}(\mathbf{x}^b) - \mathbf{c}^{a\,priori})], \quad (2)$$

where $\mathbf{A}$ denotes the averaging kernel of the retrieval, capturing the vertical sensitivity of the retrieval profiles relative to the real atmosphere, $\mathbf{c}^{a\,priori}$ is the a priori profile provided by measurement data, $\mathbf{M}(\mathbf{x}^b)$ represents a forward operator that operates on the emissions state vector and produces CO profiles that are spatially and temporally interpolated at the locations and times of the measurements, and h_v describes the vertical summation operator based on pressure weighting for computing model-equivalent column retrievals. To obtain a posteriori estimate of CO emissions, we begin by initializing the ensemble scaling factors using a multiplicative random perturbation to the prior estimates from all emission sources at the grid scale (i.e., spatially varying perturbations), sampled from based on a multivariate lognormal distribution (see the detailed description of ensemble generation in (Pendergrass et al., (2023), Section 3.2). We assume lognormal errors on the prior emissions to ensure the positivity of the solution (i.e., prevent unrealistic negative scaling factors) and to better capture the skewed tails of the emissions distribution (Maasakkers et al., 2019; Plant et al., 2022). In the next step, CHEEREIO first runs GEOS-Chem for each ensemble member over the assimilation window and then applies Eq. 2 to those ensembles in the LETKF process. This process further scales the emissions based on the observation increments (i.e., $\mathbf{y}^o - H(\mathbf{x}^b)$), and the observation and the prior error covariances (Eq.1). Note that during the construction of error covariances, a logarithmic transform of the scaling factor distributions to a normal distribution is required to satisfy the assumptions of LETKF (Hunt et al., 2007), which can be transformed back to the lognormal distribution after the LETKF process. Finally, gridded total CO emissions from all sources, including biomass burning (BB), fossil fuel, and biogenic emissions, are updated through the inversion process. Note that the analysis here will focus on regions where BB plays a dominant role in the attribution of CO emissions between May and September 2023, and where ~~the~~ BB emissions are spatially distinct from fossil fuel emissions. As a result, misattribution

of CO emissions to BB from other sources will be minor and likely falls within the uncertainty bounds of the a posteriori estimates.

To obtain an efficient performance of the inversion using TCCON data, it is important to tune the assimilation parameters with the available configuration in CHEEREIO. This was accomplished through a series of observing system simulation experiments (OSSEs), which are described in Appendices A and B. We use 24 ensemble members following previous inversion studies with the same approach (Liu et al., 2019; Pendergrass et al., 2023). Our sensitivity test with a larger ensemble size of 36 produced nearly identical posterior error estimates, with negligible improvements; thus, for saving on the computational cost of the analysis, we used the smaller ensemble size.

Before starting the inversion, we first conducted a 1-year model spin-up in 2022 (January-December) for all experiments to minimize the impact of the initial conditions on the analysis. Then, an ensemble spin-up without assimilating observations is performed, where ~~emissions are randomly~~we perturbed at each grid point emissions-based on a lognormal distribution to create an ensemble spread. We assume a lognormal standard deviation of $\sigma = 0.2$, centered around 1, that provides ensemble members scaling factors between 0.55 ~~to~~and 1.82 with 99% confidence. This perturbation level is sufficient to generate meaningful ensemble spread while avoiding unrealistically high or low values for constructing the prior error covariance. The scaling factors are assumed to be spatially correlated using an exponential decay function with a correlation distance of 500 km, while no explicit temporal correlations are imposed. We use a spin-up of about three months, comparable to the CO lifetime during summer, not only to provide a reasonable spread in the ensemble members but also to ensure the concentrations will reflect the perturbations of in emissions. ~~The choice of a lognormal distribution not only maintains the positivity of the solution (i.e., prevent an unrealistic negative scaling factors), but also better represents the uncertainty of CO emission estimates from inventories (Maasakkers et al., 2019; Plant et al., 2022).~~ At the start of the assimilation, we adjusted the ensemble members by a global multiplicative factor, making the ensemble mean equivalent to the TROPOMI and TCCON observations. This maintains a globally unbiased field of concentrations relative to the observations. Because the LETKF is sequential, it takes some time for the observations to provide sufficient information to update the emissions. To account for this lag, we use a one-month burn-in period in CHEEREIO (Pendergrass et al., 2023), for which the inversion output is discarded in postprocessing.

The performance of the inversion can be also enhanced using different LETKF parameters, such as localization and inflation parameters (Bisht et al., 2023; Miyazaki et al., 2012, 2020). For the TROPOMI and TCCON data, we use regularization factors, $\gamma_{TROPOMI} = 0.2$ and $\gamma_{TCCON} = 5$, respectively, estimated separately for each observation type through the OSSE experiments (see Appendix B). These factors scale the observation error covariances to balance the weight given to the measurements relative to the prior weight in the inversion. For the high spatial density of satellite observations, a factor of $\gamma < 1$ is required to prevent overfitting, while for sparse measurements like TCCON, a factor of $\gamma > 1$ is typically suitable to prevent underfitting of the observations. We also use an inflation factor ($\Delta = 0.08$) to compensate for a ~~rapid~~quick reduction of the ensemble spread, which may ~~otherwise result in~~prevent the inversion ~~from~~not being updated by subsequent observations. We set a localization radius of 500 km following previous CO inversion studies (Gaubert et al., 2023; Miyazaki

et al., 2015) to avoid the impact of distant observations, which may be ~~impacted~~-affected by sampling errors and spurious correlations (see Fig. S1). We generate super-observations for both TROPOMI and TCCON by aggregating measurements to the model time and grid. This aggregation helps mitigate spatiotemporal representativity errors and facilitates LETKF computations. The associated super-observation errors are calculated as described in Section 2.1. We further apply an inflation factor to the observation errors, while initially assuming an observation error correlation of 0.28, following previous studies (Chen et al., 2023; Pendergrass et al., 2023). We do not explicitly account for model transport errors in calculating the super-observations, but the influence of these errors will be captured by the error inflation in the inversion.

We conducted a series of assimilation experiments, which are listed in Table 2, to assess the impact of the use of TCCON and TROPOMI data and the choice of a priori emission inventories on the inferred CO emissions. Each experiment pairs a specific emission inventory—QFED, GBBEPx, GFAS, or CFFEPS—with either only TROPOMI observations or both TROPOMI and TCCON data in the assimilation process. Although the emissions are optimized at the grid box scale, we aggregate emissions for the following five regions (shown in Fig. 2), where there are typically significant fire emissions between May and September: North America (NA), Siberia (SI), South America (SA), Africa (AF), and South Asia and Australia (SA&A). Any emission from outside of these five regions are captured in the Rest of the World (ROW) category. The emissions from the five regions account for 90%–95% of biomass burning emissions globally.

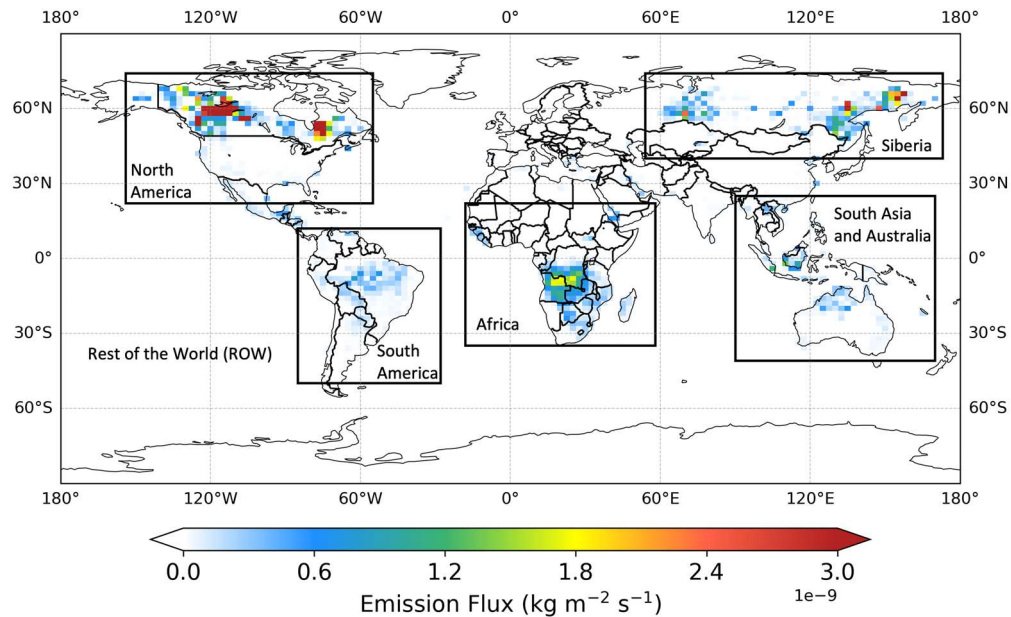


Figure 2: Major source regions of the biomass burning CO emissions used in the inversion analysis between May–September 2023, with an example of the a priori emission estimate using the GFAS bottom-up inventory. The a priori estimates for these regions from other inventories are listed in Table 3.

Table 2: Assimilation experiments with the choice of emission inventories and assimilated observations used in the global inversion between May-September 2023.

Experiment	Biomass burning emission inventory	Assimilated observations ^b
1	QFED	-
2	QFED	TROPOMI
3	QFED	TROPOMI+TCCON
4	GBBEPx	-
5	GBBEPx	TROPOMI
6	GBBEPx	TROPOMI+TCCON
7	GFAS	-
8	GFAS	TROPOMI
9	GFAS	TROPOMI+TCCON
10	CFEFS ^a	-
11	CFEFS	TROPOMI
12	CFEFS	TROPOMI+TCCON

^a The inversions based on CFEFS emission inventory in North America uses GFAS global emissions for the regions outside of North America.

^b No observations are assimilated (-), corresponding to the model a priori or control run with a particular biomass burning emission inventory.

400 4 Results and discussions

This section presents our main results by first assessing the CO emissions on a global scale (Section 4.1) and then focusing on North American emissions (Section 4.2), where the most extreme fire events took place during the study period in summer 2023. A series of experiments are conducted as listed in Table 2 for both global and regional analysis. For the global analysis (Sections 4.1.1-4.1.3), we use three biomass burning emission inventories (QFED, GBBEPx, GFAS) as priors, and
405 compare how assimilating TROPOMI satellite observations alone or jointly with TCCON ground-based measurements affects our emissions estimates. A method to measure and compare error variance and information content is presented. For evaluation, two types of independent data, including NDACC total column and surface WDCGG measurements, in addition to the same TCCON observations (non-independent) are used. Focusing on North America analysis (Sections 4.2.1-4.2.2), we include an additional regional prior emissions inventory from CFEFS (Chen et al., 2019) provided by ECCC and additional experimental
410 InSb XCO data from TCCON at ETL ~~into~~ our analysis. First, our analysis explores the spatiotemporal variability of a posteriori and a priori CO fields during extreme fire episodes. Then, as a case study at ETL, we assess the impact of assimilating additional information from experimental TCCON data at ETL, with unique retrieval characteristics, on local emissions constraints. The evaluations are performed using independent aircraft and tall tower measurements. The discussion presented

in these two sections allows us to highlight both the broad and local impacts of our approach and the specific improvements
415 achieved in areas most affected by fires.

4.1 Global analysis

4.1.1 Comparison between prior and posterior emissions

Table 3 shows the total regional BB CO emissions from the a priori bottom-up inventories and the a posteriori emission estimates (i.e., ensemble mean) obtained from the TROPOMI-only assimilation and the joint TROPOMI and TCCON (i.e.,
420 TROPOMI+TCCON) assimilation. The standard deviations shown in this table are posterior ensemble spread based on the LETKF assimilation, which is referred to as posterior uncertainty throughout the discussion in this study. The vector of posterior uncertainties has the same size as the state vector, computed for each grid cell after each assimilation window. Accordingly, the posterior uncertainty in an inversion region (i.e., regional or global scale) is the ensemble standard deviation of the total emissions. The total emissions per ensemble is obtained by summing the emissions across all grid cells in that
425 region. For the inversion period between May and September 2023, there is a large discrepancy in the a priori emissions between the three global inventories, with GFAS producing the highest global emissions of 230.3 Tg CO, which is a factor of 1.3 times GBBEPx (182.6 Tg CO) and 1.4 times QFED (164.5 Tg CO) emissions. These differences are more substantial in the regions of North America (NA) and Siberia (SI), where the boreal wildfires play an important role. In NA, the GFAS emissions are greater than the emissions from GBBEPx and QFED by a factor of 2.1 and 2.5, respectively. In SI, GFAS is a
430 factor of 3.1 and 2.1 greater than GBBEPx and QFED, respectively. Although all GFAS, GBBEPx, and QFED are based on estimates of FRP derived from satellite products (i.e., MODIS and/or VIIRS), there are many driving factors that cause the differences in the total emissions (Li et al., 2020; Liu et al., 2024), which can reach up to an order of magnitude in a finer spatiotemporal scale (Stockwell et al., 2022). In contrast, the CFFEPS emissions in NA are comparable to GFAS despite using a different approach for estimating the emissions.

435 **Table 3:** Biomass burning CO emissions estimates of a priori and a posteriori in the major source regions based on Fig. 2 for the period between May and September 2023. The a posteriori estimates are denoted by the ensemble mean and standard deviation.

Region	A priori ^a (Tg CO)				A posteriori TROPOMI-only (Tg CO)				A posteriori TROPOMI+TCCON (Tg CO)			
	QFED	GBBEPx	GFAS	CFFEPS ^b	QFED	GBBEPx	GFAS	CFFEPS	QFED	GBBEPx	GFAS	CFFEPS
North												
America	37.1	42.7	91.0	90.2	110.4 ± 19.8	112.8 ± 19.9	127.2 ± 17.2	125.6 ± 17.7	111.2 ± 15.5	115.3 ± 15.1	130.1 ± 12.9	129.3 ± 12.5
(NA)												
Africa (AF)	65.1	87.4	53.5		104.6 ± 17.3	117.2 ± 18.8	100.3 ± 17.4		103.4 ± 16.2	117.8 ± 17.7	100.8 ± 16.4	
Siberia (SI)	16.5	11.1	34.5		29.4 ± 9.8	24.1 ± 8.5	44.5 ± 12.6		29.1 ± 8.2	22.8 ± 6.9	46.4 ± 9.2	
South												
America	22.1	22.1	19.1		27.2 ± 10.3	27.4 ± 10.3	24.9 ± 10.0		28.1 ± 9.7	27.0 ± 9.4	25.0 ± 9.1	
(SA)												
South Asia & Australia	13.8	12.1	16.3		17.2 ± 6.4	15.6 ± 6.3	19.5 ± 6.8		17.2 ± 5.7	16.1 ± 5.9	20.2 ± 6.1	
(SA&A)												
Rest of the World	9.8	7.2	5.8		14.1 ± 3.0	12.5 ± 3.0	12.2 ± 3.2		18.9 ± 2.7	12.6 ± 2.2	16.2 ± 2.7	
(ROW)^c												
Global	164.5	182.6	230.3		302.9 ± 66.6	309.6 ± 66.8	328.6 ± 67.1		307.9 ± 58.0	311.6 ± 57.1	338.7 ± 56.4	

^a A priori Fossil fuel and biofuel emissions are provided by CEDS inventory and biogenic emissions are obtained from MEGAN version 2.1 in all the cases.

^b The global inversions with CFFEPS BB emissions uses GFAS global BB emissions for the regions outside of North America.

^c The source of CO in the rest of the world also includes the oxidation of methane and non-methane hydrocarbons.

440 After assimilating the TROPOMI data, the a posteriori emissions suggest a large increase in global emissions from the a priori, with global a posteriori emissions of all inventories in close agreement with each other. We estimate global fire emissions of 303±67 Tg CO, 310±67 Tg CO, and 329±67 Tg CO for QFED, GBBEPx, and GFAS, respectively. In North America, the a posteriori emission estimates are also consistent, with estimates of 126±18 Tg CO, 127±17 Tg CO, 113±20 Tg CO, and 110±20 Tg CO for CFFEPS, GFAS, GBBEPx, and QFED, respectively. For most other regions, the inferred emissions all agree with within the a posteriori uncertainty except for Siberia, where the a posteriori emission estimates are 45±13 Tg CO, 24±9 Tg CO, and 29±10 Tg CO for GFAS, GBBEPx, and QFED, respectively. The discrepancy between the a posteriori Siberian emissions could be because the assimilation did not ingest TROPOMI observations poleward of 60°N, and there are large emission sources poleward of 60°N, as can be seen in Fig. 4. Note that although the emissions beyond 60° may not be directly corrected from local observations at the current assimilation step, they can still be updated using observations between 60°S-60°N and through model transport and cycling of the assimilation that propagate information globally. The overall

agreement between the a posteriori emission estimates obtained with the different a priori emissions suggest that TROPOMI provides sufficient information to constrain the regional emission estimates.

The joint TCCON and TROPOMI inversion produces a posteriori emission estimates that agree to within 5% of the a posteriori emissions from the TROPOMI-only inversion, but these differences vary among the source regions. At the global scale, the a posteriori estimates remain within 1σ uncertainty, implying that the increment from ~~the~~ joint inversion closely matches that from the TROPOMI-only ~~assimilation~~inversion. Previous inversion studies with CO, and CO₂ measurements showed that combining satellite total column observations with surface in situ measurements (Byrne et al., 2020; Kim et al., 2024; Wang et al., 2018) could benefit from the complementarity between the two types of measurements and provide posterior fluxes that are better informed by measurements. We will discuss later in this section the spatiotemporal distributions of the a posteriori emissions in comparison with the a priori ~~between-across all-the~~ inventories. Table 3 shows that the joint inversion reduces the uncertainties in the posterior relative to the inversion using only the TROPOMI data. We find a global uncertainty reduction of nearly 15% in all the a posteriori emission estimates when including the standard TCCON XCO in the inversion. This reduction is likely due to the additional temporal information and higher accuracy provided by TCCON, compared to TROPOMI. The reduction of uncertainty varies between the source regions. In the Northern Hemisphere extratropics, where most ~~of-the~~ TCCON stations are located, the uncertainty reduction ~~reaches~~increases to 29% in NA with CFFEPS emissions, followed by Siberia with GFAS emissions (26% reduction ~~of uncertainty~~). On the other hand, for the tropics and subtropics in the Southern Hemisphere, where there are fewer TCCON stations, we find smaller reductions in uncertainties (between 5% and 10%). Goudar et al. (2023) investigated the uncertainties in estimating CO emissions from isolated fires using TROPOMI assimilation. They reported that these estimated uncertainties primarily arise from errors due to spatial under-sampling of the CO field by TROPOMI observations and errors due to assumptions about the temporal variability of the emissions. Although we did not examine individual fire events, the lower overall uncertainty from the joint TCCON and TROPOMI assimilation suggests an improved handling of the spatial under-sampling error in TROPOMI-only assimilation, which is reflected in the uncertainty estimates. This improvement could be particularly important, as it shows the isolated effect of adding TCCON data to the inversions, when other factors such as a priori emissions and their errors were kept fixed between the joint and TROPOMI-only assimilations.

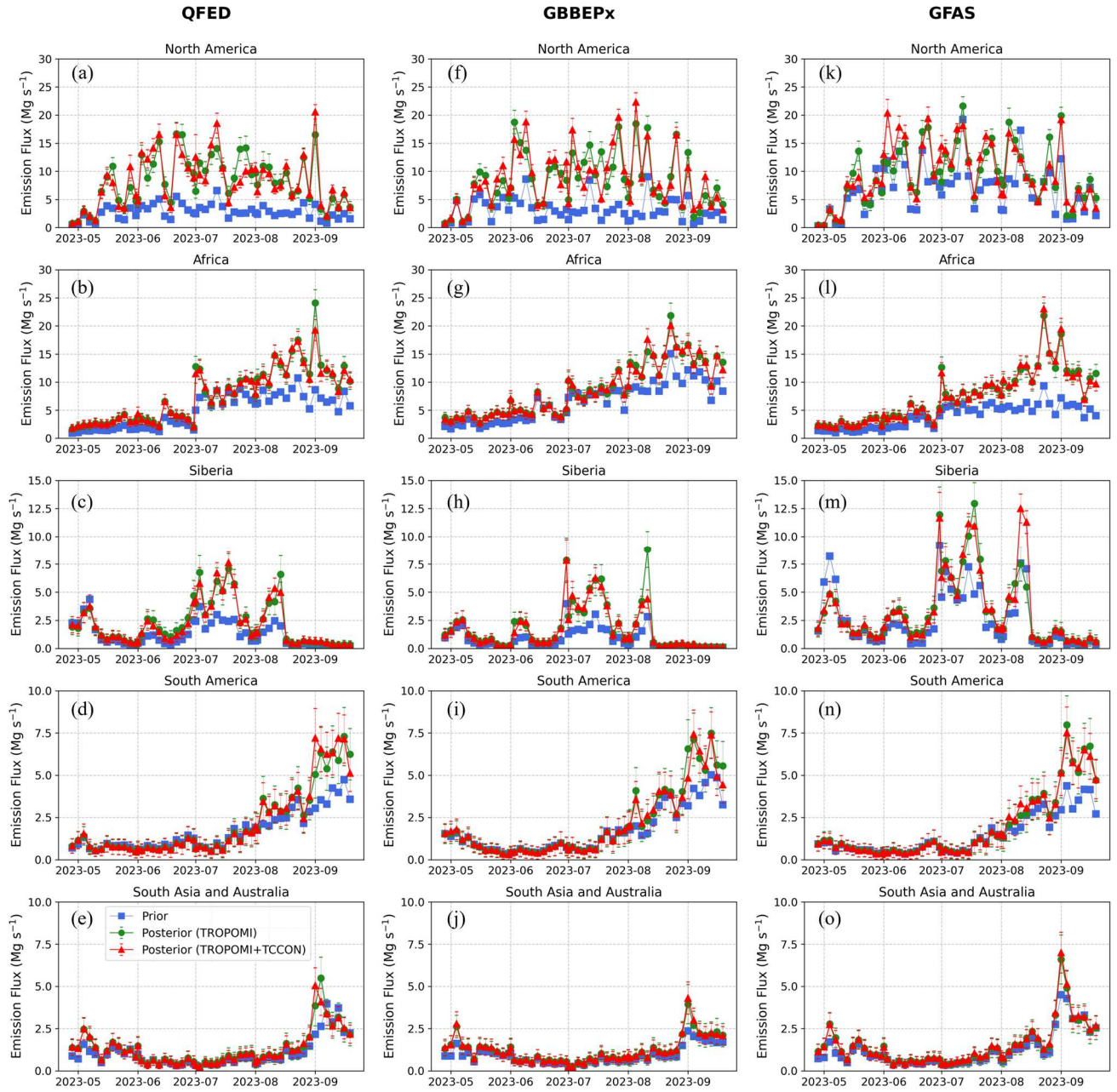
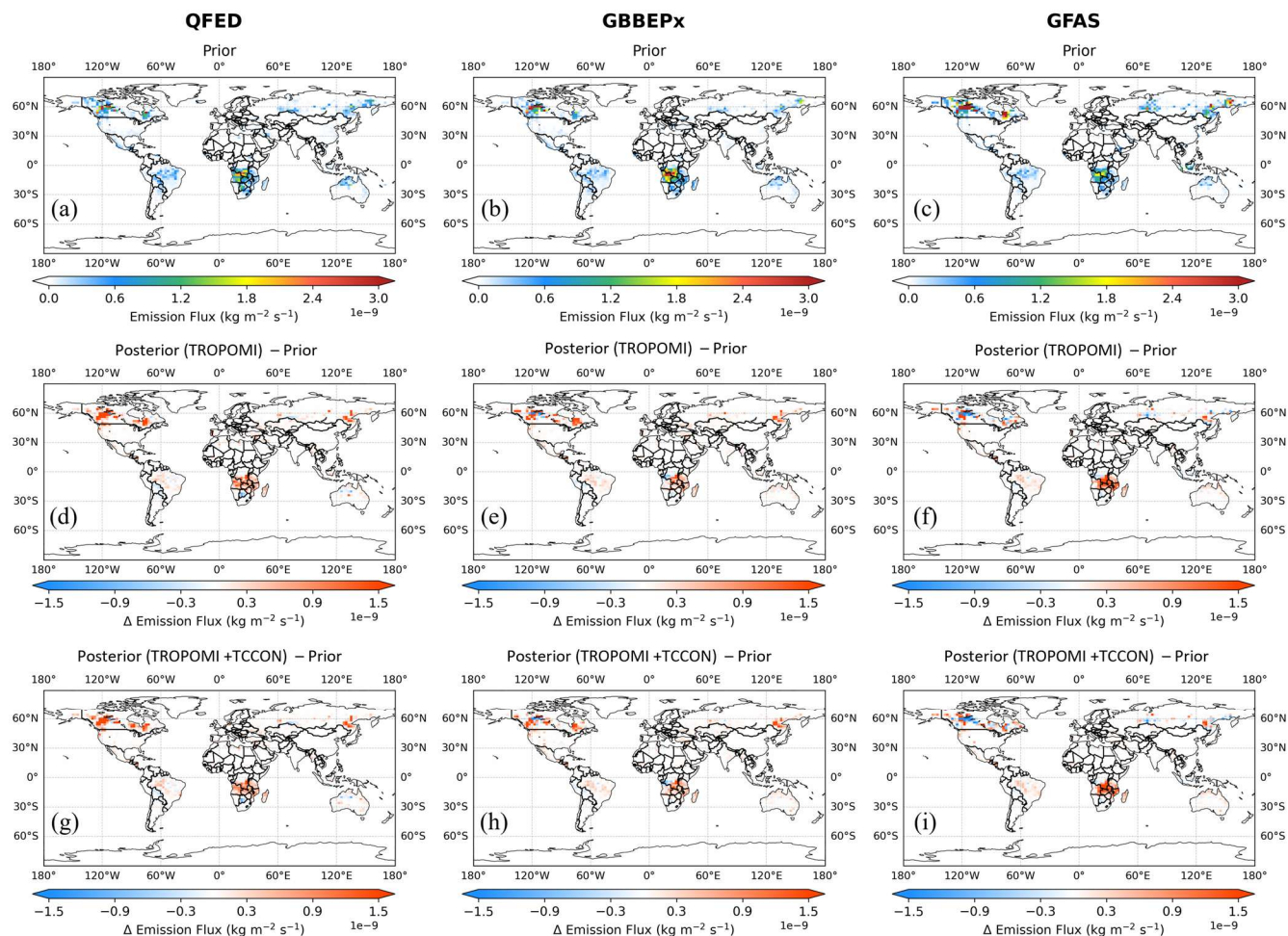


Figure 3: Comparison of the temporal variability of the CO emissions estimate for the priors (blue), posteriors using TROPOMI-only assimilations (green), and posteriors using joint TROPOMI and TCCON assimilations (red), among the three global biomass burning inventories, including (a-c) QFED, (f-j) GBBEPx, and (k-o) GFAS used as the priors, and for the major inversion source regions as shown in Fig. 2.



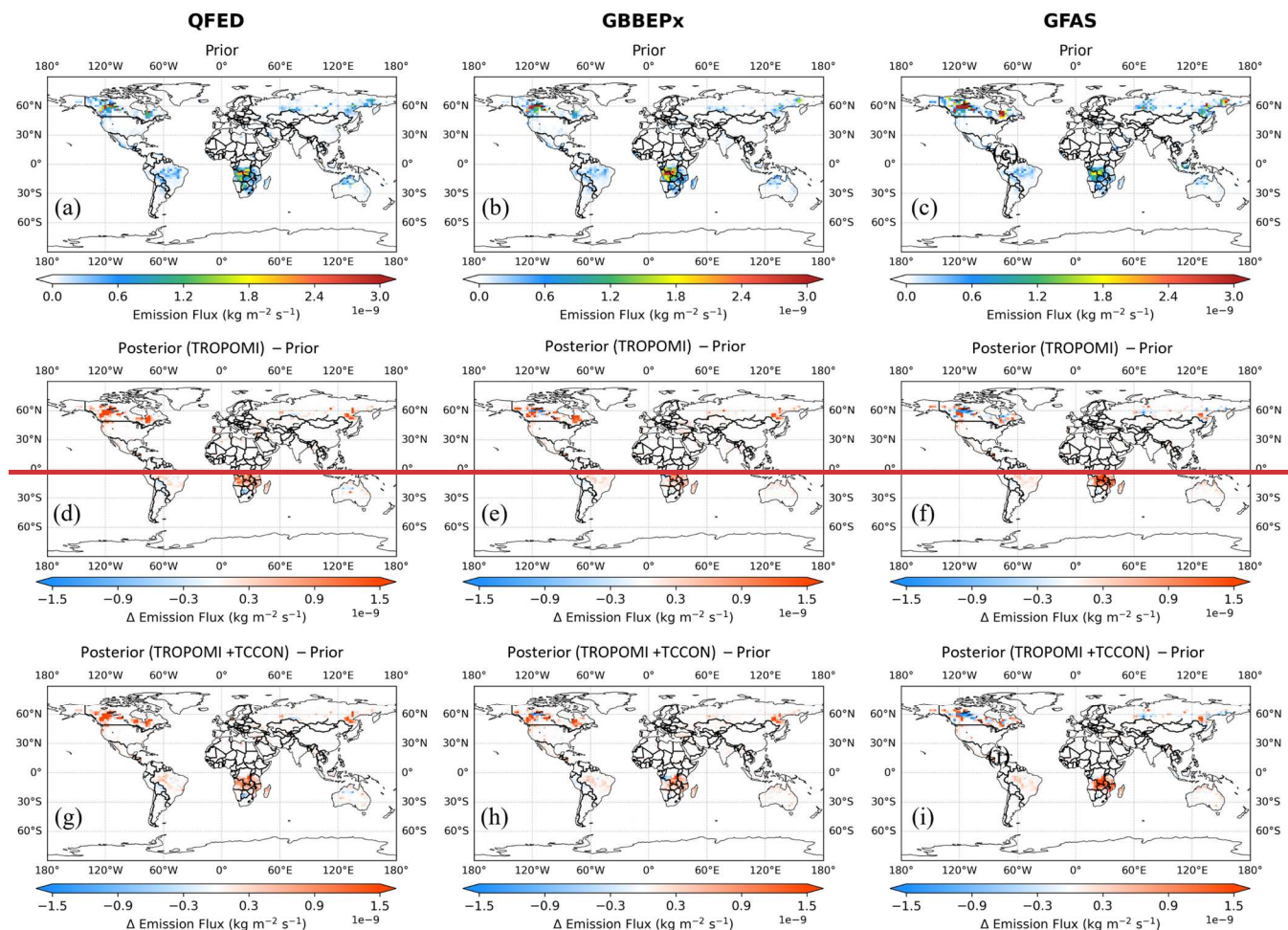


Figure 4: Comparison of the spatial distribution of the time-averaged biomass burning CO emissions in the (a-c) a priori, (d-f) a posteriori – a priori using TROPOMI-only assimilation, and (g-i) a posteriori – a priori using joint TROPOMI and TCCON assimilation, and among three global biomass burning inventories.

To better understand the influence of assimilating TCCON XCO together with the TROPOMI XCO, we examine the temporal and spatial variability of the estimated emissions. Fig. 3 shows a priori emissions in blue, a posteriori emissions from the TROPOMI-only assimilation in green, and the a posteriori emissions from the joint inversion in red between May and September. In North America, QFED shows only slight variations in the a priori emissions (Fig. 3a). GBBEPx (Fig. 3f) shows some degree of improvement over QFED relative to TROPOMI posterior during a few fire episodes. For GFAS, however, this improvement over QFED (Fig. 3k) is significant, so that the GFAS prior exhibits reasonable agreement in both magnitude and temporal variability with the a posteriori emissions from the TROPOMI-only assimilation. The posteriors from different priors, but with the same set of assimilated observations, show overall stronger temporal variability compared to their respective priors. For instance, in North America, while the QFED and GBBEPx priors are relatively flat compared to the GFAS prior, all their posteriors exhibit enhanced and more consistent variability (Fig. 3a,f,k). Still, differences to some extent remain

noticeable between these posterior time series, likely driven by differences in the spatial and temporal distribution of the priors, and their interaction with observational constraints through model transport and mixing. Comparison of the spatial distribution of the a priori emissions (Fig. 4a-c) indicates that two main regions of boreal wildfires in (eastern) Quebec and (western) Alberta and British Columbia, Canada, correspond to the large differences in regional emissions among the inventories, although their overall global spatial distributions are similar. The a posteriori – a priori emissions from the TROPOMI-only inversions confirm the underestimate of CO emissions in QFED and GBBEPx, in those regions in Canada (Fig. 4d-f). GFAS, unlike the other two inventories, has significantly larger emissions from wildfires across Canada, so that they are comparable with the magnitude of the a posteriori emissions in that region (see Table 3 for a comparison of total emissions and Fig. S2 in the supplements for the separate a posteriori emission maps associated with Fig. 4).

In Siberia, the a posteriori emissions provide an enhancement on the a priori in a few episodes in July and August (Fig. 3c,h,m), indicating an overall low level of emissions from all inventories. GFAS, followed by QFED, has not only higher emissions in the same locations as GBBEPx, but also has a greater area of BB emissions. However, the a posteriori emissions suggest that the GFAS inventory required larger adjustments in the central and western part of SI, suggesting errors in the spatial allocation of the a posteriori fire emissions in Siberia. In Africa and South America, the emissions enhancement occurs in late July through the end of the inversion period in September, which are shown in Fig. 3b,g,l,d,i,n. For these regions, GBBEPx, followed by QFED, provide a closer estimate to the a posteriori than GFAS, while the posterior emissions from the TROPOMI-only and joint inversions remain nearly identical (Fig. 4d-f versus 4g-i). This could be because there are very few TCCON measurements in the Southern Hemisphere near that region. In South Asia and Australia, the emissions enhancement occurs only in two episodes in early May and early September, which remain nearly consistent among the inventories. Finally, comparing the regional a posteriori emissions from the TROPOMI-only and joint inversions (Fig. 3k,l,m,n,o), we find different temporal variability of the updated emissions, which are distinctive for several wildfire episodes. Additionally, we observed improvements in the correlation between analogous posterior timeseries across different BB inventories. For example, the temporal correlation between the QFED and GFAS posterior emissions in North America increased from $r = 0.85$ in the TROPOMI-only assimilation to $r = 0.90$ in the joint TROPOMI and TCCON assimilation. These findings suggest that the higher temporal resolution of the TCCON measurements provide additional constraints on temporal variability of the emissions in the inversion, which is consistent with findings from previous studies on CO₂ inverse modelling (Byrne et al., 2020, 2024; Chevallier et al., 2011).

Overall, according to the discussion above and the results from the Observing System Simulation Experiments (OSSEs) demonstrated in Appendices A and B, we find that although TCCON alone may not significantly constrain spatiotemporal variability in the major inversion regions—likely due to the limited number of measurement sites—it is still clear that adding TCCON to TROPOMI in the joint inversion reduces the posterior uncertainty estimates everywhere compared to the TROPOMI-only inversion. We found that the reduction of the uncertainty by adding TCCON measurements becomes more significant during high BB emission episodes or wildfire events. Later, in Section 4.1.3, we evaluate the a priori together with the a posteriori from both TROPOMI-only and joint inversions using independent measurements.

4.1.2 Error variance reduction and the information content

We evaluate the performance of the inversion for constraining BB CO emissions by quantifying the information content provided by the TROPOMI and TCCON data. To achieve this, we use two approaches: (i) computing the reduction of uncertainty in the model space and (ii) computing the degree of freedom for signal (DOFS) in the ensemble subspace. Using these methods will quantify the information provided by the two observing systems individually. In the first approach, we compute the a priori and a posteriori error variance for each grid point, which is obtained as part of the solution for LETKF processing in CHEEREIO. The reduction of error variance can be used as a metric for evaluating the inversion performance (Feng et al., 2009); as such, a greater error variance reduction at a grid point indicates that more reliable information from observing system is available to constrain emissions for that grid. Accordingly, we define a normalized error reduction (ε) for each grid point as follows:

$$\varepsilon_i = 1 - \frac{(\sigma_i^a)^2}{(\sigma_i^b)^2}, \quad (3)$$

where $(\sigma_i^a)^2$ and $(\sigma_i^b)^2$ denote the error variance of the a posteriori and the a priori, respectively, for the i^{th} emissions in the state vector. ε_i varies between 0 to 1, with greater values indicating a higher reduction of a priori uncertainties. Although the LETKF method approximates the a posteriori uncertainty due to the reduced rank representation associated with the limited number of ensembles used to construct the error covariances (Livings et al., 2008), it still provides useful information with which to evaluate the analyses.

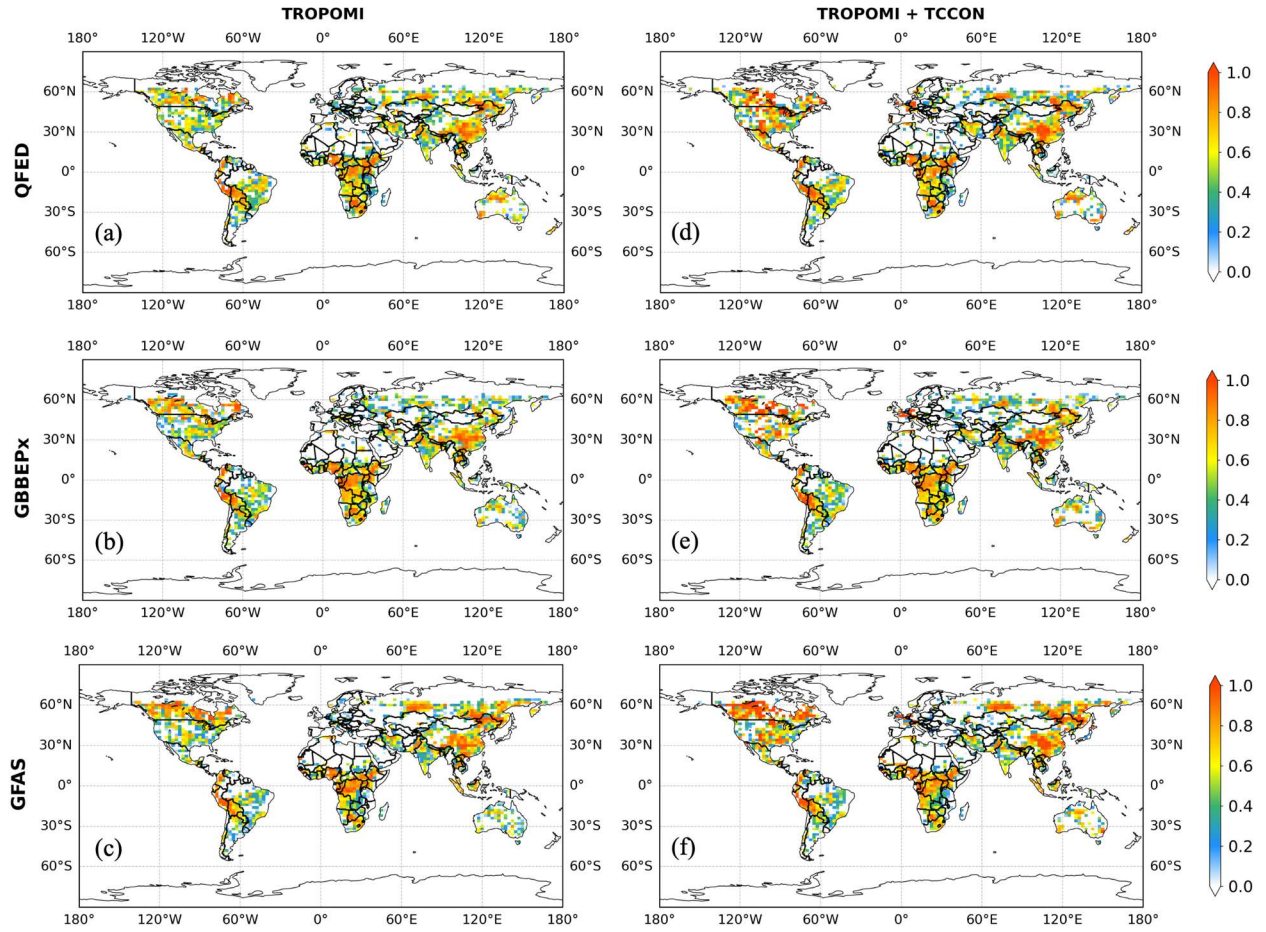


Figure 5: Comparison of the time-averaged (May-September 2023) a posteriori error variance reduction, ε , of (a-c) TROPOMI-only assimilation and (d-f) joint TROPOMI and TCCON assimilation for three global biomass burning inventories used as a prior.

Fig. 5 shows the error variance reduction, ε_i , for the a posteriori emissions based on the three different global BB inventories (i.e., QFED, GBBEPx, and GFAS) using TROPOMI-only measurements (Fig. 5a-c) and using joint TROPOMI and TCCON measurements (Fig. 5d-f). The greater reduction in error variances implies higher confidence in the posterior estimate at those locations. We find greater reduction of error variance in the joint inversion compared to the inversion with TROPOMI-only data, primarily in NA and in the vicinity of TCCON stations. In fact, the reductions correlate with the weight of observations compared to the model a priori, so that increasing the weight of observations with respect to the model a priori could result in higher reductions of the error variances. For example, with QFED as the a priori emissions, the rate of reduction is greater in boreal Canada, central and southern part of the United States, western Europe, eastern Asia, Siberia, and Australia. This indicates where TCCON provides additional information to further constrain the emissions based on the QFED BB inventory. In addition, comparing the cases with different inventories suggests that there could be differences in error variance reduction due to the model a priori. Accordingly, the slightly greater reduction with GFAS, compared to QFED or GBBEPx,

560 is likely due to its higher spatiotemporal variability, which enables the inversion to better exploit the information from observations.

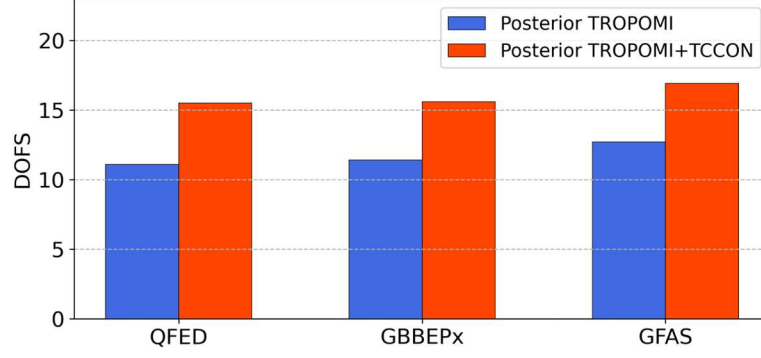


Figure 6: Computed degree of freedom for signal in ensemble subspace (DOFS_k) associated to the TROPOMI-only (blue) and joint TROPOMI and TCCON (red) inversions for three global biomass burning inventories (QFED, GBBEPx, GFAS) used as a priori.

565 In the second approach, following Zupanski et al. (2007), we compute the degree of freedom for signal (DOFS) approximated for the ensemble-based assimilation method. The DOFS , as defined by Rodgers (2000), quantifies the number of pieces of independent information in an observing system toward constraining the state vector of dimension n (also equivalent to the total number of grids in the model). It is defined as

$$\text{DOFS}_n = \text{tr}(\mathbf{I}_{n \times n} - \mathbf{P}_{n \times n}^a (\mathbf{P}_{n \times n}^b)^{-1}) = \text{tr}(\mathbf{A}_{n \times n}), \quad (4)$$

where $\mathbf{P}_{n \times n}^a$ and $\mathbf{P}_{n \times n}^b$ are analysis and background error covariances, $\mathbf{I}_{n \times n}$ denotes the identity matrix, and $\mathbf{A}_{n \times n}$ represents the averaging kernel matrix. To compute $\mathbf{A}_{n \times n}$ and then DOFS_n , a Jacobian matrix must be constructed in full rank, requiring extensive computational cost (e.g., Varon et al., 2022). In an ensemble-based approach, computing the Jacobian matrix in the state space is impractical since the limited number of ensembles are not sufficient to describe full rank error covariances. However, those quantities can be approximated for the ensemble subspace with reduced rank error covariance matrices, so that the information provided by the observation system is measured relative to the maximum independent pieces of information determined by the ensembles size, k . Therefore, the DOFS_k is defined as,

$$\text{DOFS}_k = \text{tr}((\mathbf{I}_{k \times k} + \mathbf{C}_{k \times k})^{-1} \mathbf{C}_{k \times k}) = \text{tr}(\mathbf{A}_{k \times k}), \quad (5)$$

where $\mathbf{C}_{k \times k}$ denotes the symmetric information matrix, $\mathbf{I}_{k \times k}$ is the identity matrix, and $\mathbf{A}_{k \times k}$ represents an equivalent averaging kernel or influence matrix, all obtained in the ensemble subspace. The derivation of Eq. (5) started from Eq. (4) is described in Zupanski et al. (2007) and Zupanski (2005). Subsequently, we can compute the information matrix, $\mathbf{C}_{k \times k}$, either within the LETKF calculation of CHEEREIO or as a postprocessing step if all the outputs from the ensemble members and the control run are already stored. Accordingly, every element of matrix $\mathbf{C}$, are computed as

$$\mathbf{C}_{ij} = \mathbf{z}_i^T \mathbf{z}_j, \quad (6)$$

where $\mathbf{z}$ is the vector of dimension m (the number of observations) and is defined as

$$\mathbf{z}_i(m) = \mathbf{R}_{m \times m}^{-\frac{1}{2}} (H_m(\mathbf{x}_i) - H_m(\mathbf{x})), \quad (7)$$

where the observation operator H_m is applied to the perturbed, $\mathbf{x}_i$, and unperturbed state vector $\mathbf{x}$, and weighted by the inverse of square root of observation error covariance $\mathbf{R}_{m \times m}$. Migliorini (2013) used the same method in the square root filter, where the forecast error covariance matrix is approximated by the sample covariance matrix, which is produced by the forecast of each ensemble member. Once matrix $\mathbf{C}$ is constructed, one can use Eq. (5) to obtain $\mathbf{A}_{k \times k}$ and the DOFS _{k} for the ensemble subspace. Note that using this approach, there are at most $k - 1$ independent pieces of information for the entire assimilation period. Thus, with $k = 24$ in this work, the computed DOFS may vary between 0 and 23, mainly depending on the characteristics of the assimilated observations, such as their density and error statistics. Although this method does not produce DOFS in the state space of the emissions (Žagar et al., 2016), it enables a straightforward comparison of the information content across different experiments.

The computed DOFS for different inversions are shown in Fig. 6. Adding TCCON data to the inversion increases the DOFS from the TROPOMI-only inversion for all the cases. These values increase from 11.1, 11.4, and 12.7 for the TROPOMI-only inversions to 15.5, 15.6, and 16.9 for the inversions with TROPOMI and TCCON data using QFED, GBBEPx, and GFAS emissions, respectively. The higher DOFS from GFAS compared to the other BB inventories is also in agreement with its higher reduction of uncertainty in Fig. 5. This likely implies that the difference between the perturbed and unperturbed forecast of the state vector, which defines the elements of matrix $\mathbf{C}$, correlates with the spatiotemporal variability of the prior emissions. Thus, GFAS prior, with greater variability than the other priors (e.g., Fig. 3a,f,k and Fig. 4a-c), may result in higher DOFS.

4.1.3 Evaluation using ground-based observations

We evaluate the inversion against TCCON, NDACC total column retrievals, and in situ WDCGG measurements to better understand the constraint from each measurement type used in the inversion. NDACC and WDCGG are independent data while TCCON are the same data as those used in the assimilation, so they are not independent for evaluating the joint inversions. Table 1 shows the measurement sites with their geographical information (latitude, longitude, and altitude above sea level). First, we evaluate the results against all the TCCON measurements from May to September 2023. In Fig. 7, the model is evaluated against hourly averaged TCCON data, with the a priori shown in blue and the a posteriori in red from either the TROPOMI-only inversions (in the left) or the joint TROPOMI and TCCON inversions (in the right). The statistics indicate that the inversions significantly improve on the prior for all the cases. Examination of the a priori models shows that the GBBEPx simulation slightly improves on the QFED simulation, with a coefficient of determination of $R^2 = 0.31$ compared to $R^2 = 0.27$ with QFED. The GFAS simulation has the highest a priori correlation ($R^2 = 0.54$), resulting in the best a posteriori agreement with TCCON, with $R^2 = 0.82$ and $R^2 = 0.87$ for the TROPOMI-only and TROPOMI+TCCON assimilation, respectively. The time series plots in Fig. 7 show the improvement of GBBEPx on QFED, with better agreement between the simulation and the TCCON measurements when there are wildfire enhancements of XCO. The GFAS simulation shows a significant improvement during those peaks, indicating that the GFAS simulation better captures the time variability in the

measurements. The evaluation of the joint inversion using TROPOMI and TCCON XCO data against the TCCON measurements shows a further improvement of the inversions; both the slope and the R^2 are closer to 1.0 than the results from the TROPOMI-only assimilation. This is expected, because the same TCCON data are used in the inversion and the evaluation (Fig. 7d,h,l).

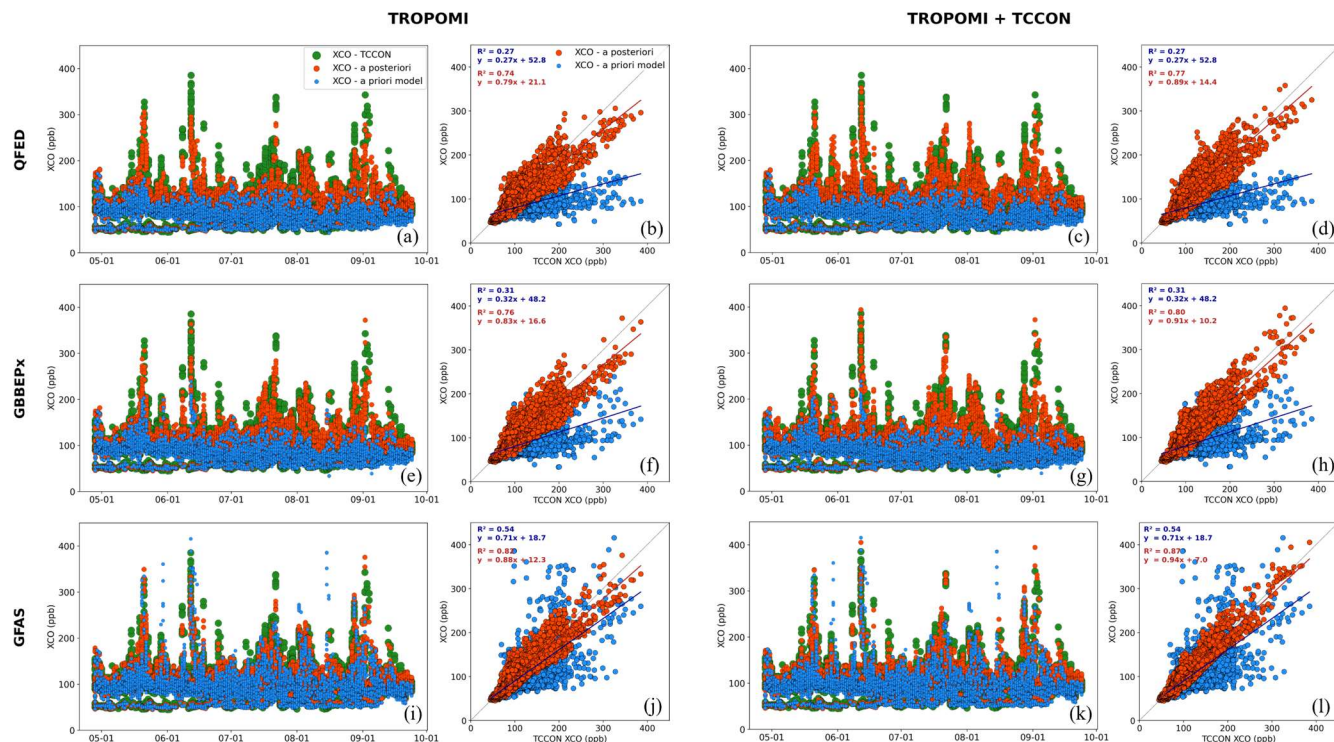


Figure 7: Evaluation of the model a priori (blue) and a posteriori from TROPOMI-only assimilation (red) against TCCON measurements (green) for all sites together using a time series and scatter plots based on (a, b) QFED, (e, f) GBBEPx, and (i, j) GFAS biomass burning emissions (left side). A similar evaluation of the model a priori against TCCON, however but with a posteriori from joint TROPOMI and TCCON assimilation (red) using (c, d) QFED, (g, h) GBBEPx, and (k, l) GFAS as the prior biomass burning emissions estimates (right side).

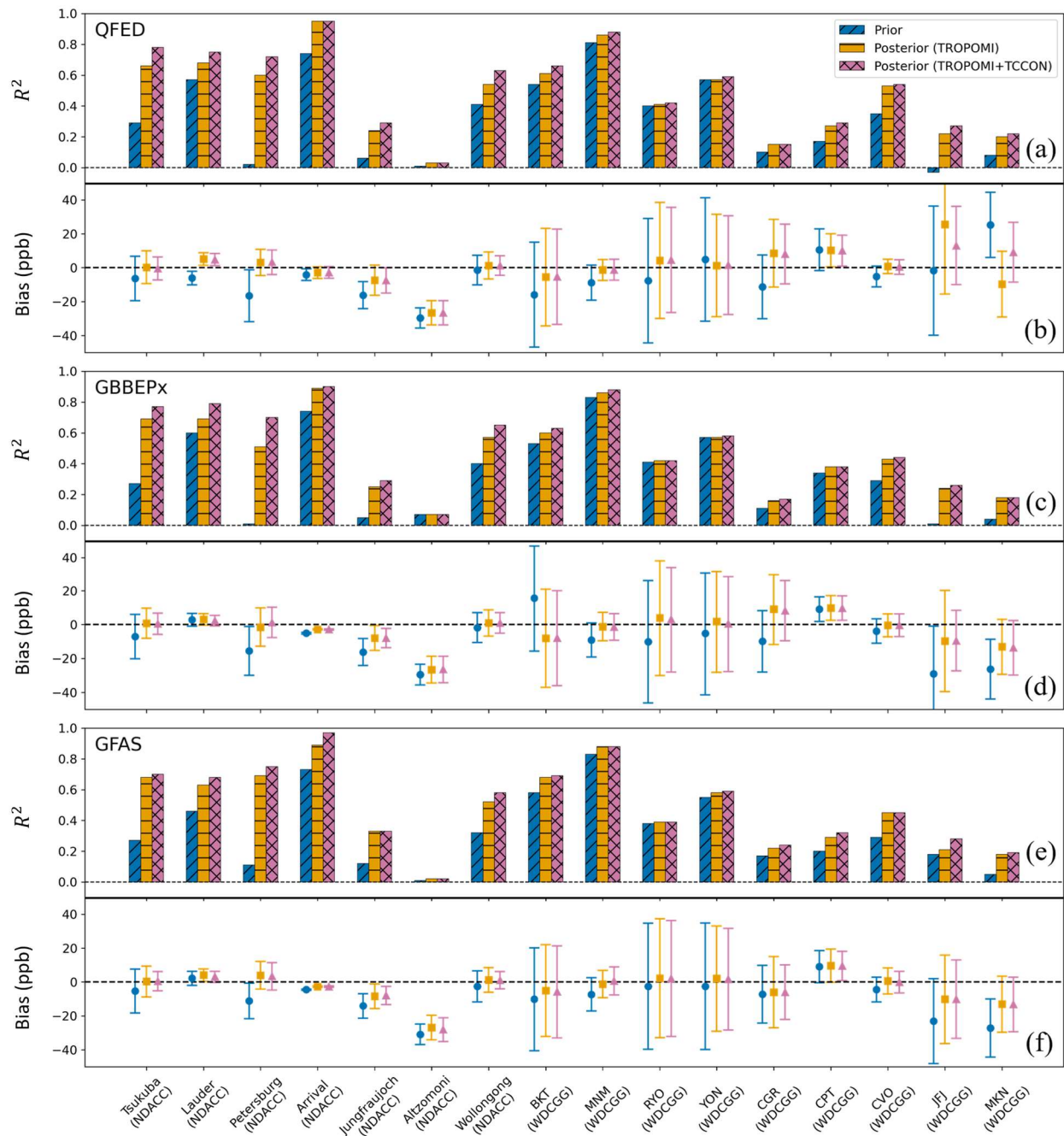


Figure 8: Evaluation of the time-averaged model a priori (blue), a posteriori using TROPOMI-only assimilation (orange), and a posteriori using joint TROPOMI and TCCON observations (red) against independent NDACC and in situ measurements between May-September 2023. Each panel is associated with a particular prior biomass burning emissions ~~from bottom-up inventories~~, including (a, b) QFED, (c, d) GBBEPx, and (e, f) GFAS. The top row of each panel shows the coefficient of determination (R^2) of the model using the prior or assimilation using the posterior with respect to the measurements, whereas the bottom row represents the mean bias and the standard deviation of the model or assimilation with respect to the measurements. Tsukuba, Lauder, and Wollongong are collocated stations for both TCCON and NDACC.

To assess the impact of TCCON on the performance of the inversion more objectively, we also compare the inversion results with independent NDACC column measurements and surface in situ measurements in Fig. 8. The R^2 , mean bias, and standard deviation relative to the measurements are shown for the model a priori (QFED, GBBEPx, and GFAS) in blue, the a posteriori from the TROPOMI-only assimilation in orange, and the a posteriori from the joint TROPOMI and TCCON assimilation (i.e., TROPOMI+TCCON) in red. We find a higher R^2 for both a posteriori estimates (i.e., TROPOMI-only and TROPOMI+TCCON) relative to the a priori estimates in almost all the cases, while there is an additional increase in correlation for the joint inversion compared to the TROPOMI-only inversion. The a posteriori from the TROPOMI-only assimilation provides a small reduction of the mean bias and standard deviation, and by adding TCCON to the assimilation, there is a further reduction at several sites (Tsukuba, St. Petersburg, MNM, RYO, YON). The added improvement in the posterior XCO obtained by adding the TCCON data to the inversion differs between sites. For the NDACC sites collocated or downwind of TCCON sites, such as Tsukuba, Lauder, Wollongong (collocated stations), and St. Petersburg (~2000 km downwind of European stations), the R^2 increases more by adding TCCON to the inversion. The evaluation at the Arrival Heights NDACC station located in a remote area in Antarctica, far from both the TROPOMI and TCCON assimilated observations, shows an improvement in the a posteriori that suggests that the assimilation improves global background concentrations of CO. However, at the Altzomoni NDACC site, located about 75 km southeast of Mexico City and almost 2 km higher in altitude, we found little improvement after the assimilation. It is likely that the local topography cannot be captured in our global model, which has a $2^\circ \times 2.5^\circ$ spatial resolution. In addition, the local ambient atmospheric conditions, such as stability and humidity in this region, causes most of the fire emissions to stay within the boundary layer, and neither the model nor the assimilation are-is capable of capturing such effects (Sha et al., 2021).

For the surface in situ measurements, we also find an increase of-in the correlation between the a priori and the a posteriori estimates using the TROPOMI+TCCON assimilation, but with a smaller improvement than we-seeobserved at the NDACC stations, with the exception of the JFJ and MKN stations, where there is a larger improvement. The relatively smaller improvement compared to NDACC may primarily be attributed to the fact that the surface sites have a larger representativeness error given our $2^\circ \times 2.5^\circ$ grid resolution, which limits the ability of the assimilation to significantly correct for them. For example, for-at the CGR station, the local atmosphere is influenced by a land-sea wind regime that cannot be resolved by the relatively coarse grid resolution of the model. In addition, the vertical sensitivity of the TCCON and NDACC data based on their averaging kernels may partly impact these evaluations.

Similar to the evaluation against NDACC, we find slight improvements in the mean bias and standard deviation at WDCGG surface stations. In most of-the-cases, adding TCCON to the assimilation reduces the error standard deviations (i.e., posterior-measurement errors), while the mean bias remains almost identical to the-bias-that of the TROPOMI-only inversion. Comparison among the three BB cases shown in Fig. 8 indicates that, for an-equivalent a priori or a posteriori emissions estimates (e.g., for TROPOMI-only), their statistics are not significantly different at most measurement sites.

4.2 North America analysis

4.2.1 Assimilation performance for constraining boreal wildfires emissions

Our posterior emissions from the inversions indicated that North America has the highest level of BB emissions and contributed one-third of the global total in summer 2023. The emissions primarily came from the boreal forest across Canada, which were poorly estimated by the bottom-up emissions inventories, with a 31%–67% underestimation in the a priori relative to the a posteriori from both TROPOMI-only and joint inversions in this region during the study time-period (see Section 4.1 and Table 3). Thus, we take a closer look at the spatial and temporal characteristics of the XCO over North America to better understand the localized and episodic behavior of the fire emissions with respect to the different inventories. In addition to the three global BB emissions inventories discussed in the previous sections, here, we also include a regional bottom-up emissions inventory in North America from CFFEPS (Chen et al., 2019) provided by ECCC to compare with those global emissions. Note that for our global simulation with CFFEPS emissions in North America, we use GFAS for the global emissions which are replaced by CFFEPS in North America.

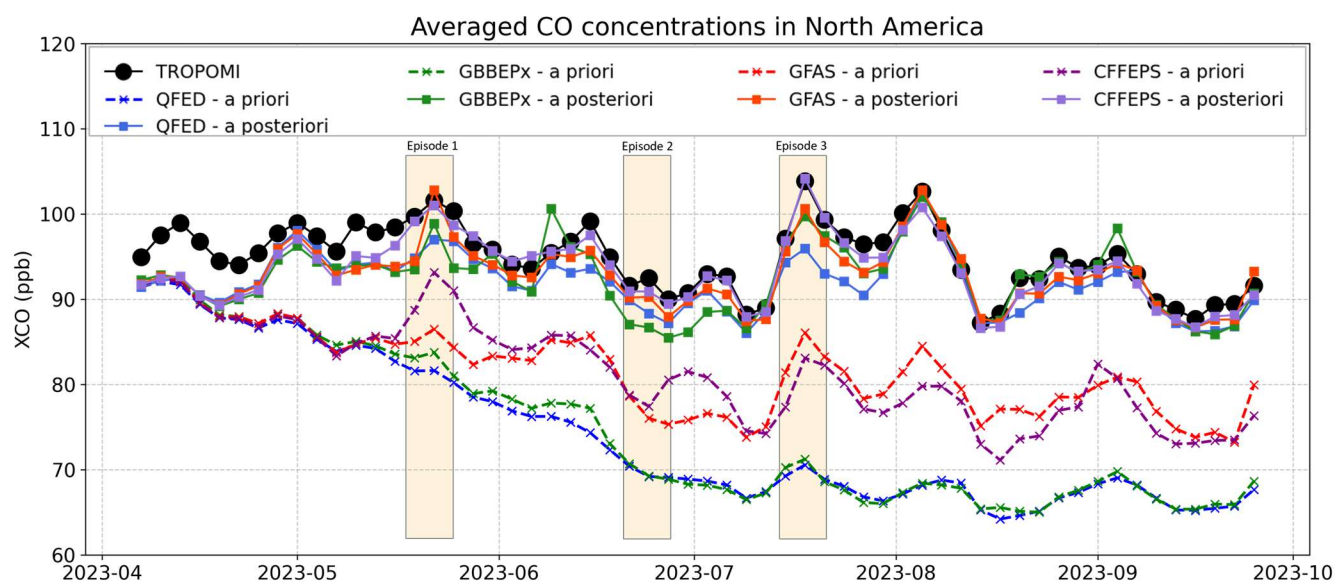


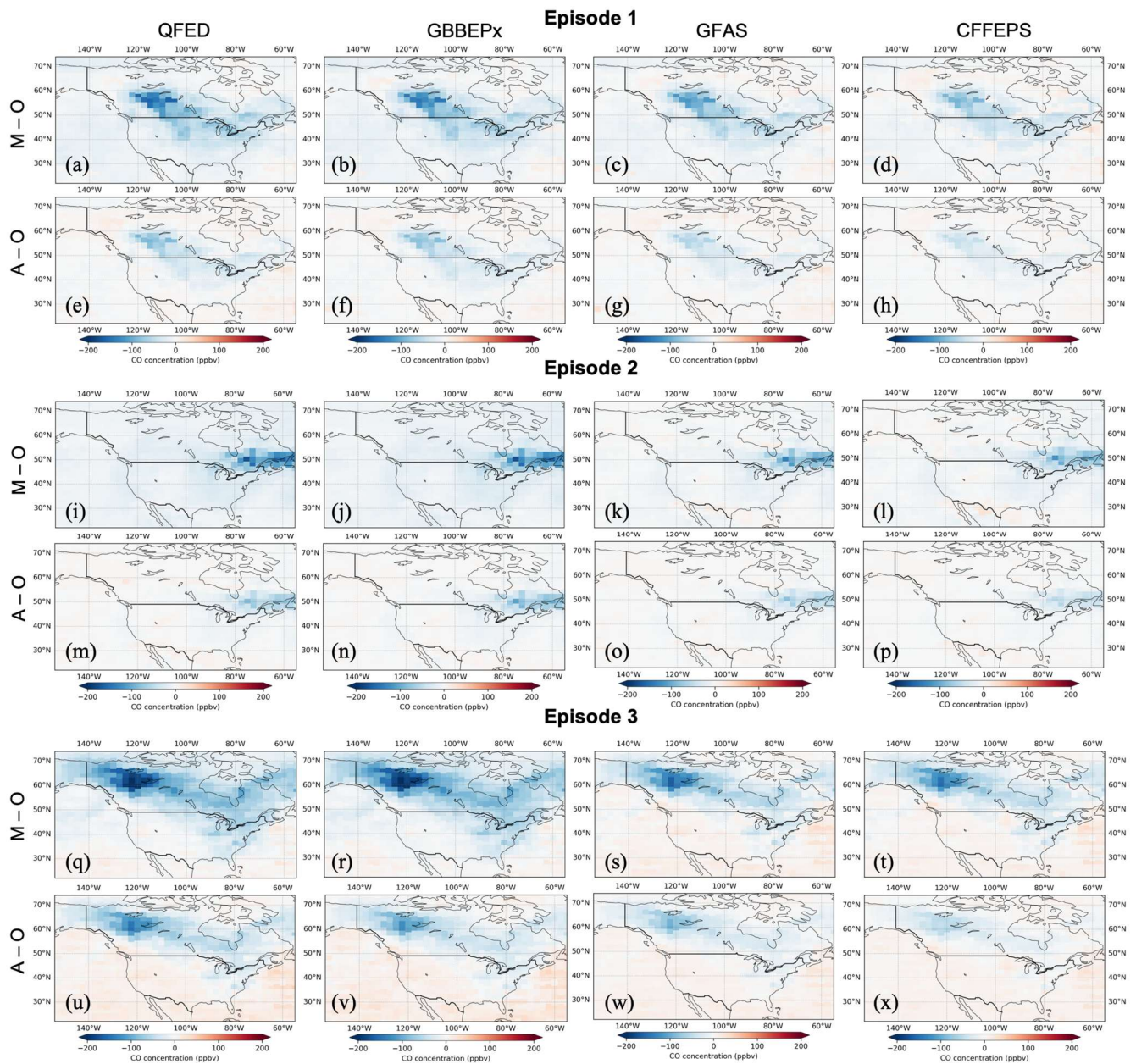
Figure 9: Evaluation of the domain-averaged CO concentrations (XCO) of the a priori model (dashed lines) and a posteriori using TROPOMI assimilations (solid lines) between-for four different inventories in North America, including QFED (blue), GBBEPx (green), GFAS (red), and CFFEPS (purple), against TROPOMI XCO measurements (black). The a priori model refers to a model forecast using prior emissions and the a posteriori is equivalent to the ensemble mean from LETKF. Three extreme wildfire episodes across boreal regions are chosen for comparison of-between the assimilation results using different inventories and for comparison of the assimilation with the a priori model.

We first focus on the temporal variability of the domain-averaged XCO in North America from the a priori model and a posteriori assimilation-estimate using TROPOMI inversion with the four inventories. In Fig. 9, the model a priori with-for each emission inventory is shown in dashed lines with an 'x' marker, the a posteriori is shown with solid lines with a square marker, and the TROPOMI measurements themselves are indicated by the black line with circles. Fig. 9 shows that the a priori

using QFED and GBBEPx emissions have similar XCO in the entire period, except a slightly higher level of XCO with GBBEPx during May and June. On the other hand, the a priori XCO estimates with GFAS and CFFEPS are both greater ~~than QFED and GBBEPx~~ in magnitude (~~i.e.,~~ ~10 ppb higher) and more variable ~~than those with QFED and GBBEPx~~. Although, the model-estimated CO ~~has shows~~ similar trends between CFFEPS and GFAS, the two inventories produce XCO with different temporal variability and both underestimate the XCO observed by TROPOMI. In fact, during the large emissions episodes from wildfires in mid-May and late June, CFFEPS has higher emissions and better captures the variability in TROPOMI XCO, ~~while-whereas~~ from July to September, they produce comparable levels and variability of CO₂ with slightly higher CO for GFAS at the peaks in late July and early August.

Nevertheless, the inversion using TROPOMI suggests that all the a posteriori emission estimates are a significant improvement from the a priori after about a month, and show reasonable agreement with the temporal variability of the TROPOMI measurements. The a posteriori XCO also suggests that the seasonal variability that is usually characterized by decreasing CO in summer, due to the higher rate of oxidation with OH radical, is balanced by the higher rate of BB CO emissions estimated in the inversions, resulting in an almost uniform XCO during the summer-fall 2023. Despite their poorer a priori estimates, QFED and GBBEPx provides a posteriori XCO that agrees with TROPOMI measurements better than the a priori of GFAS and CFFEPS. However, comparing the XCO between the four inversions suggests that CFFEPS, followed by GFAS, perform better at the XCO peaks, and thus can better capture the variability in the TROPOMI measurements.

We looked closely at three extreme wildfire episodes that occurred across the Canadian boreal forest at different times and regions in summer 2023 to better examine the spatial characteristics of the a priori and a posteriori estimates at the time of the fires. As shown in Fig. 9, the first episode covers five days of large wildfires in Alberta between 19th and 23rd of May, the second episode ~~of extreme wildfires~~ occurred in Nova Scotia and Quebec between the 22nd and 26th of June, and the third episode was in British Columbia and the Northwest Territories between the 17th and 19th of July. To evaluate the inversion results from these events, we compare in Fig. 10 the model a priori (M) and a posteriori analysis (A) with the TROPOMI observations to obtain analysis minus observations (A – O) and model minus observations (M – O) differences for the four emissions inventories during the three extreme wildfire episodes in North America. A comparison of M – O between the different inventories for all the episodes reveals that CFFEPS, followed by GFAS, has a smaller underestimation of CO concentrations compared to QFED and GBBEPx. Although the reduction of this bias occurred over a large domain, including downwind of the emissions, the reduction is more significant in the vicinity of the wildfire emissions. Our results comparing different inventories are consistent with our findings in Fig. 9, which show that CO concentrations from the a priori are underestimated due to the lower emissions in the inventories (Fig. 4), in the same order as observed here. We find similar improvements in our a posteriori analysis between the different inventories, in which A – O exhibits lower bias with ~~CFFEPS~~CFFEPS, followed by GFAS, in comparison with QFED and GBBEPx. Although the CFFEPS a posteriori XCO is significantly closer to the TROPOMI observation, the QFED a posteriori still shows a slight improvement on the CFFEPS a priori, indicating the larger impact of assimilating TROPOMI observations compared to providing a better prior.



720 **Figure 10:** Comparison of the difference between model (M) or assimilation (A) and TROPOMI XCO observations (O) (i.e., $M - O$ or $A - O$), for three extreme wildfire episodes across boreal regions. Episode 1: May 19–23; episode 2: June 22–26; episode 3: July 17–19. Assimilation is based on TROPOMI observations and uses different inventories in North America as the prior CO emission estimates. The model and assimilation fields were transformed using the TROPOMI a priori profiles and averaging kernels.

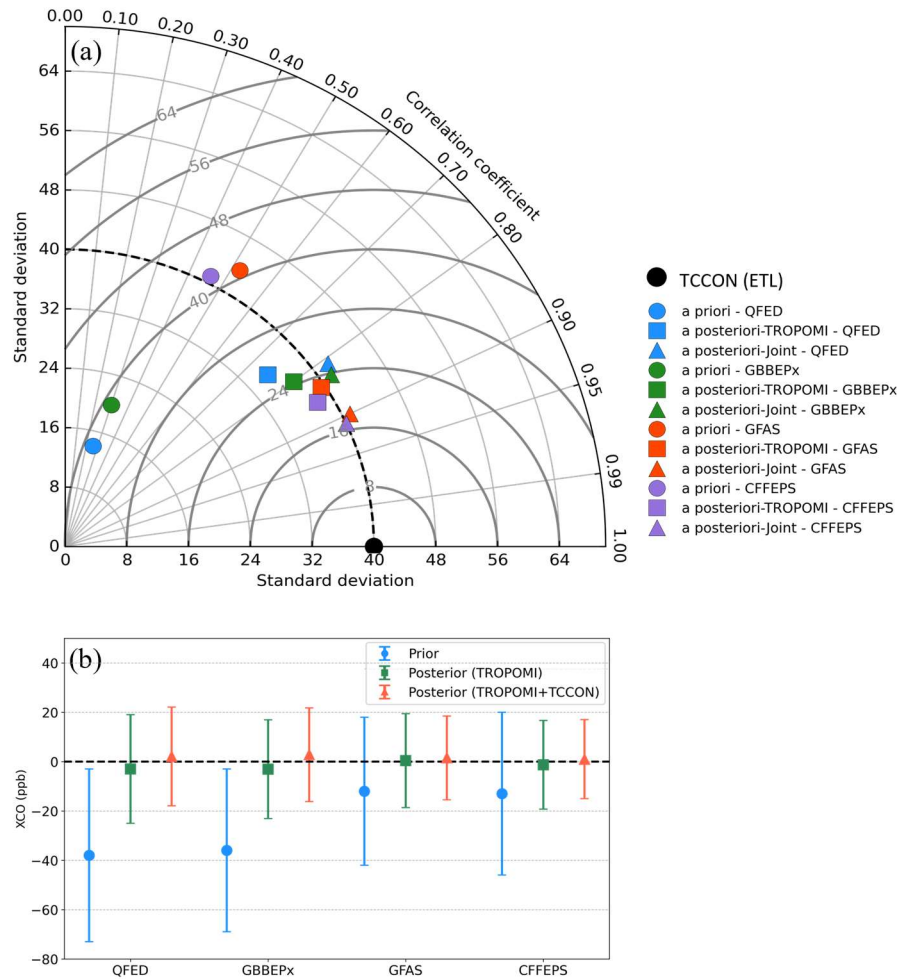


Figure 11: Taylor diagram for evaluation of the assimilation and a priori model using four different biomass burning inventories against TCCON XCO measurements at East Trout Lake (ETL) between May and September 2023. (a) Assimilation is performed using TROPOMI-only (square) and joint TROPOMI and TCCON (triangle) data, while their correlations and standard deviations are compared with the model a priori (coloured circles) and TCCON XCO measurements (black circle). (b) Evaluation of the mean bias and error of the priors (blue), assimilation using TROPOMI-only data (green), and assimilation using a joint TROPOMI and TCCON data (red).

We also evaluated the a priori and a posteriori XCO with TCCON data from ETL for the entire simulation period between May and September 2023. Fig. 11a presents a Taylor diagram comparing the standard deviation and correlation of the a priori or a posteriori against TCCON measurements at ETL, and Fig. 11b shows the mean bias and standard deviation of the prior/posterior—measurement residuals. The a priori XCO are shown in circles, the a posteriori from the TROPOMI-only assimilation in squares, and the a posteriori from the joint TROPOMI+TCCON assimilation in triangles. It shows that the a priori estimates with QFED and GBBEPx have not only low correlation with the measurements, but also have low variability. The a priori estimates with CFFEPS and GFAS improve on the correlation but more significantly on the variability, in addition to the mean bias (Fig. 11b) that is reduced by more than a factor of 2. The a posteriori for all cases provides significant

improvements on the a priori by increasing the correlation and lowering the mean bias and standard deviation, resulting in closer estimates to the measurements. Among all the a posteriori cases, the joint TROPOMI and TCCON inversion has a noticeable level of improvement with increased correlation and slightly smaller mean bias and standard deviation, in addition to adjusting the variability towards the measurement's variability. A comparison among the inventories suggests that the joint inversion using CFFEPS provides the highest correlation, lowest standard deviation, and nearly unbiased estimates with matching variability matching with the measurements. The best agreement between the joint inversion and the TCCON measurements at ETL is with CFFEPS, followed by GFAS, then GBBEPx, and QFED.

4.2.2 Implications for vertical sensitivities in the inversion

In this section, we examine the potential use of the experimental TCCON XCO product from the mid-infrared (InSb) detector available at ETL for the inversions of BB CO emissions. Since the measurements provide us with an independent set of XCO₂ with distinct averaging kernels, we take those as separate pieces of information into our joint TROPOMI and TCCON inversion. Similar to the standard TCCON, we assume uncorrelated errors for the InSb data, however, the effect of possible error correlations can be approximated and taken into account by re-tuning the regularization factor. The InSb CO measurements are processed in a similar way to produce a separate set of hourly gridded TCCON InSb data. Thus, when these data are added to the joint inversion, the size of the observation vector and observation error covariance increases accordingly. The vertical profile of the averaging kernels in the InSb CO product (Fig. 1d) has higher sensitivity to the surface and lower troposphere, and lower sensitivity to higher altitudes compared with the standard XCO. Thus, we aim to understand the added benefit of assimilating these data for constraining CO emissions in the inversion. To achieve this, we conducted inversions using joint TROPOMI and TCCON data that incorporate three variations of the TCCON product, including the standard XCO, the InSb XCO, and the combined InSb and standard TCCON product. The inversions show nearly identical improvements between the posteriors and the prior when using the CFFEPS inventory, indicating a less than 1% discrepancy of total BB emissions in North America. However, there is a noticeable difference in their spatial distributions, especially in the wildfire hotspots in British Columbia, Alberta, and Quebec (see Fig. S3 in the supplements).

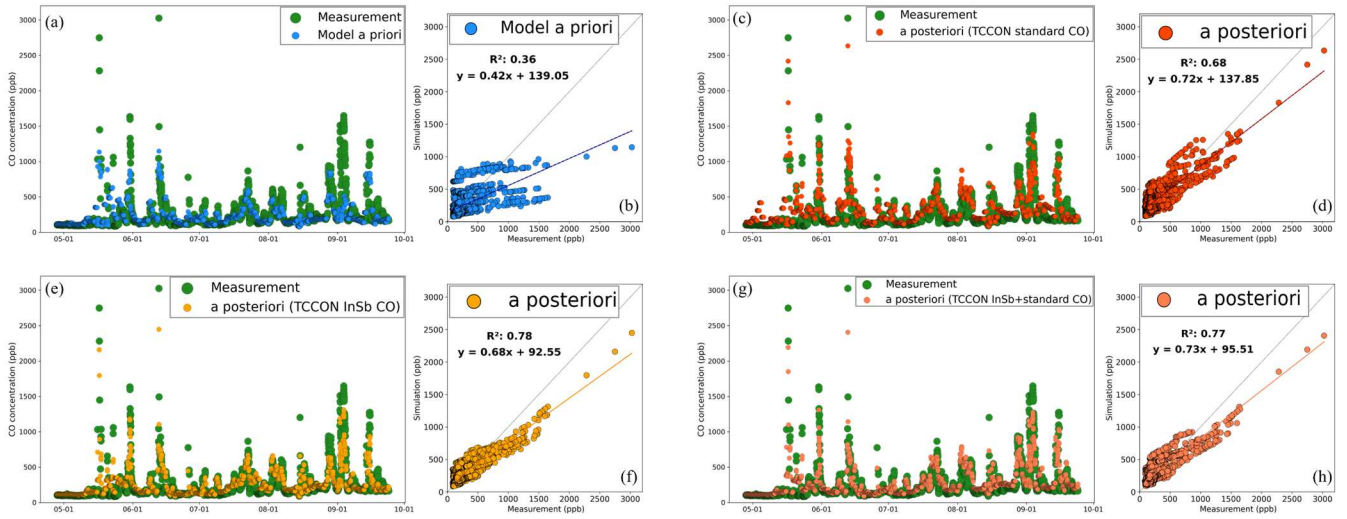


Figure 12: Evaluation of the assimilation and a priori model using the CFFEPS biomass burning inventory against tall tower measurements at ~60 m above sea level at East Trout Lake (ETL). (a-b) a priori model (blue), (c-d) assimilation using joint TROPOMI and TCCON standard TCCON data (red), (e-f) assimilation using joint TROPOMI and InSb TCCON data (yellow), and (g-h) assimilation using joint TROPOMI and the standard and InSb TCCON data (orange) are compared with tall tower measurements (green).

To evaluate the performance of these inversions, we use two independent in situ datasets: tall tower measurements from Environment and Climate Change Canada (Chen et al., 2014) and aircraft profiles from the National Oceanic and Atmospheric Administration (McKain et al., 2024; https://gml.noaa.gov/aftp/data/trace_gases/co/pfp/aircraft/) at ETL. We compare the tall tower measurements with the a priori (in-blue) in Fig. 12a,b and the a posteriori (in-red) from the joint TROPOMI and TCCON assimilation using standard TCCON in Fig. 12-c,d, the a posteriori (in-yellow) from the joint assimilation using the TROPOMI and InSb TCCON data in Fig. 12e,f, and the a posteriori (in-orange) from the combined TROPOMI plus the standard and InSb TCCON data in Fig. 12g,h. We find a significant improvement in all the inversion cases compared to the a priori, especially at the concentration peaks, resulting in an increase in R^2 and the slope of the regression.

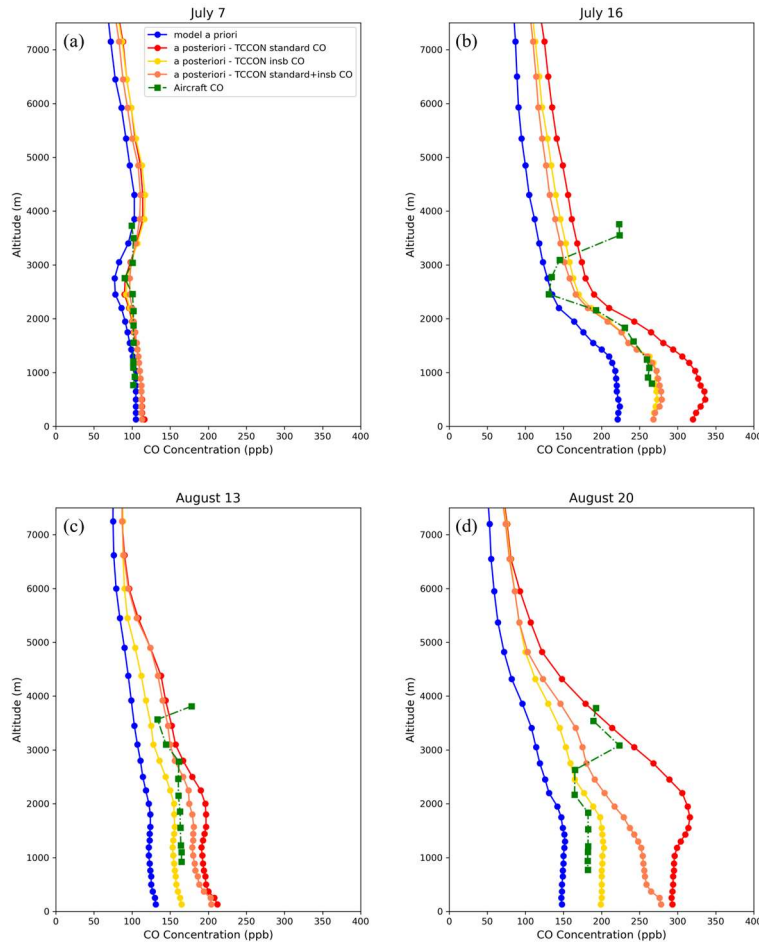


Figure 13: Evaluation of vertical profiles from the joint inversion using TCCON standard CO (red), InSb CO (yellow), combined standard and InSb CO (orange), and the a priori model (blue) against aircraft measurement profiles from NOAA at ETL (green squares) on (a) July 7, (b) July 16, (c) August 13, and (d) August 20, 2023.

The evaluation of the inversions shows that the assimilation with TROPOMI and the InSb XCO data improves the correlation from the case with standard TCCON XCO (from $R^2 = 0.68$ to $R^2 = 0.78$), although the slope of the regression has slightly decreased (from $slope = 0.72$ to $slope = 0.68$). However, the inversion using the combined standard and InSb TCCON products improves on both correlation and the slope of the regression ($slope = 0.73$, $R^2 = 0.77$) with respect to the inversion using the standard TCCON. This suggests that the inversion using the InSb XCO better captures the variability of CO near the surface, likely associated with the greater sensitivity of these data to lower altitudes, which improves the sensitivity of the data to surface emissions in localized regions with short-range transport. The slightly lower slope is likely due to the greater level of underestimation of CO at the peak concentrations compared to the standard TCCON inversion. This might be because the lower sensitivity of the InSb XCO to the mid-troposphere than the standard CO, could reduce the measurement sensitivity to CO plumes at higher altitudes or in the background, which is normally captured through longer range transport. However, the combined standard TCCON and InSb CO assimilation captures the variability slightly better than the standard

790 TCCON CO assimilation, with improved correlations relative to independent data, and also provides us with more representative estimates of the background CO. Thus, adding the InSb XCO dataset potentially benefits the inversion by providing a better constraint on the surface BB CO emissions.

Furthermore, an evaluation of the vertical profiles of CO from the a priori and a posteriori simulations against aircraft in situ measurements by NOAA at ETL in Fig. 13 shows that there is a consistent improvement with the joint inversion using different variations of TCCON data (the inversion using standard TCCON CO is shown in red, whereas the inversion using the InSb TCCON CO is shown in yellow), and the combined standard and InSb TCCON CO (in orange). Note that through the inversion process, we update only CO emissions, without directly updating concentrations. The results indicate that replacing the standard TCCON with the InSb product improves the agreement with the measurements at lower altitudes (1-2.5 km), while at higher altitudes (2.5-3.5 km) the standard TCCON assimilation performs slightly better. Despite the fact that a perfect constraint on vertical profiles cannot be obtained by assimilating only total column measurements, due to the limited vertical sensitivity, using both the standard and InSb CO data together in the assimilation maintains a balanced and reasonable agreement with the independent measurements at both lower and higher altitudes. This suggests that using all the TCCON standard and InSb CO measurements in the inversion provides an improved constraint on the fire plumes at a broader range of altitudes. This is likely associated with uniformly larger sensitivities with altitude compared to the inversion using each of the TCCON XCO datasets individually. Note that we have not found a similar level of improvement from adding the ETL InSb XCO dataset to our inversions when we evaluate against aircraft in situ data at the Park Falls TCCON station (not shown), which is about 1700 km to the southeast of ETL. This suggests that, although adding the InSb XCO to the inversion benefits the inversion results, it has a more local effect, and may not provide a substantial additional constraint on the regional/global scale. Therefore, providing the InSb product at other TCCON locations is recommended for a better constraint on the emissions on larger scales.

5 Summary and conclusions

We used total column measurements from the TROPOMI satellite and the TCCON ground network to infer CO biomass burning emissions during the extreme North American fire season between May and September 2023. Using the CHEEREIO toolkit, we optimized CO emissions globally at a $2^\circ \times 2.5^\circ$ grid resolution every 3 days. One objective of this work is to better understand the influence of the TCCON measurements in providing additional constraints for quantifying CO emissions through a joint TCCON and TROPOMI inversion. Despite the limited spatial coverage, TCCON has substantially more observations in time with high accuracy on column-averaged dry mole fraction measurements. This motivates the evaluation of the joint inversion in comparison with the TROPOMI-only inversion to constrain emissions from localized and episodic wildfires. A second objective is to evaluate the global QFED, GBBEPx, and GFAS a priori BB emission inventories, as well as the regional North American CFFEPS emissions, and to assess their impact on the inversion analyses.

All of the inversion results indicate that the priors significantly underestimate the BB CO emissions. Based only on TROPOMI observations, the global posterior emission estimates for QFED, GBBEPx, and GFAS are 302.9 ± 67 , 309.6 ± 67 , and 328.6 ± 67 Tg CO, compared to prior estimates of 164.5, 182.6, and 230.3 Tg CO, respectively. For North America, the posterior emissions for QFED, GBBEPx, GFAS, and CFFEPS are greater than the priors by a factor of 3.0, 2.6, 1.4, and 1.4, respectively. Adding TCCON through a joint inversion with TROPOMI makes little difference to the global total and regional estimates ($< 5\%$), but it improves the temporal variation as well as the spatial distribution in the BB hotspots. Furthermore, we found that the joint inversion reduces the uncertainty of the posterior in all major inversion regions, but with different magnitudes, reaching near 30% over North America. The spatial distribution shows that the uncertainty reductions are larger in proximity and upwind of the TCCON measurement sites. Our evaluation of the information content in the ensemble subspace also indicates that the joint TCCON and TROPOMI inversion increase the DOFS by ~~between 33~~—39%, depending on the prior inventory. However, the additional constraints provided by the TCCON data correlate with the spatial density of the TCCON sites (more sites in the Northern Hemisphere), such that a greater benefit is obtained in North America and Siberia than in Africa and South America.

The evaluation of the results against the TCCON measurements and independent NDACC column and in situ surface measurements obtained from WDCGG reveals that the TROPOMI-only inversion primarily improves on the biases while the joint inversion further increases the correlations. The joint inversion can better capture the temporal variability of the measurements, resulting in a more accurate estimate at the peak concentrations during the extreme wildfire events. The statistics also reveal that the standard deviations and the mean errors of the differences between the assimilation and measurements are lower in the joint inversion in comparison with the TROPOMI-only inversion, providing us with more reliable estimates of atmospheric CO. These improvements due to the joint inversion are not consistent throughout all the measurement sites we used for evaluation; there is stronger agreement at the NDACC and in situ sites that are located in close proximity to the TCCON measurements used in the inversion. Thus, the spatial distribution of TCCON is also a factor driving the improvements in the inversion.

Our comparison using different inventories suggests that, although they all provide similar spatial distributions of BB CO emissions, their magnitude and temporal variability can be different. For example, our evaluations against TROPOMI in North America suggest that GBBEPx has slightly higher and GFAS has significantly higher emissions than QFED. CFFEPS has an overall similar level of emissions to GFAS, but provides enhanced temporal variability that is in better agreement with the TROPOMI measurements. However, the inversion posteriors with all inventories indicate a significant improvement on the priors, such that the posterior obtained with QFED, which has the lowest prior emissions, provides posterior CO that is in better agreement with the measurements than the priors based on GFAS or CFFEPS, which had higher emissions. However, the impact of better prior emissions is not negligible; for instance, the posterior with CFFEPS produces CO which agrees better with observations. In fact, CFFEPS, in addition to lowering spatial biases of the inversion results (Assimilation – Observation), more closely captures the temporal variability of the measurements with the addition of the TCCON data in the inversion.

Finally, we investigated the potential use of the experimental TCCON InSb CO measurements at ETL in the inversion.

855 The experimental TCCON XCO provides greater sensitivity than the standard TCCON XCO to the surface. The evaluation of the inversion results against independent measurements suggests that replacing the TCCON standard CO with the InSb CO results in posterior CO that correlates better with surface measurements, although the mean bias slightly increases. This suggests that using InSb CO increases the sensitivity to the local surface emissions, but may degrade the sensitivity to transported plumes or the background CO. However, the inversion~~s~~ with both the standard and InSb CO data of TCCON
860 improve~~s~~ on the standard CO inversion, provides higher correlations and lower mean biases relative to the surface measurements. The evaluation against aircraft data also emphasizes the potential benefit of using InSb CO measurements in the inversion for constraining surface emissions, as it improves the agreements with observations at lower altitudes.

Our results using TCCON measurements in a joint inversion with TROPOMI data suggest that increasing the spatiotemporal density of observations allows the assimilation to constrain CO emissions at finer scales, providing useful
865 information for a more reliable estimation of local and episodic wildfires. Furthermore, we showed the benefits of the TCCON measurements² accuracy and temporal density when they are used jointly with TROPOMI in the inversion. However, assimilating TCCON data alone is not sufficient to fully constrain the spatial context. Integrating more observations from ground networks, such as NDACC, COCCON, and in situ observations (e.g., Schuldt et al., 2024), with satellite observations offers the potential to greatly enhance the performance of inversion analyses for quantifying fire emissions.

870 **Appendix A: OSSEs with joint TCCON and TROPOMI inversion**

Observing System Simulation Experiments (OSSEs) are widely used to evaluate the behaviour of atmospheric inversion or assimilation systems by using simulated observations under ~~an~~-idealized conditions~~s~~. OSSEs allow us to explore how individual observation datasets or the underlying system setup (e.g., assimilation parameters) can enhance the overall performance of the system (Lahoz and Schneider, 2014; Bocquet et al., 2015; Abida et al., 2017; Voshtani et al., 2023). Here,
875 we conduct twin experiments (Ghil and Mo, 1991)~~;~~ to evaluate the implementation of the inversion system. Specifically, we assess the potential utility of TCCON observations for quantifying CO emissions. Additionally, we use the OSSEs to estimate assimilation parameters and error statistics of the background and observations, using diagnostics such as chi-square tests (Ménard and Chang, 2000; Tang et al., 2024) or statistics based on observation-minus-forecast (OmF) (Miyazaki et al., 2020; Voshtani et al., 2022).

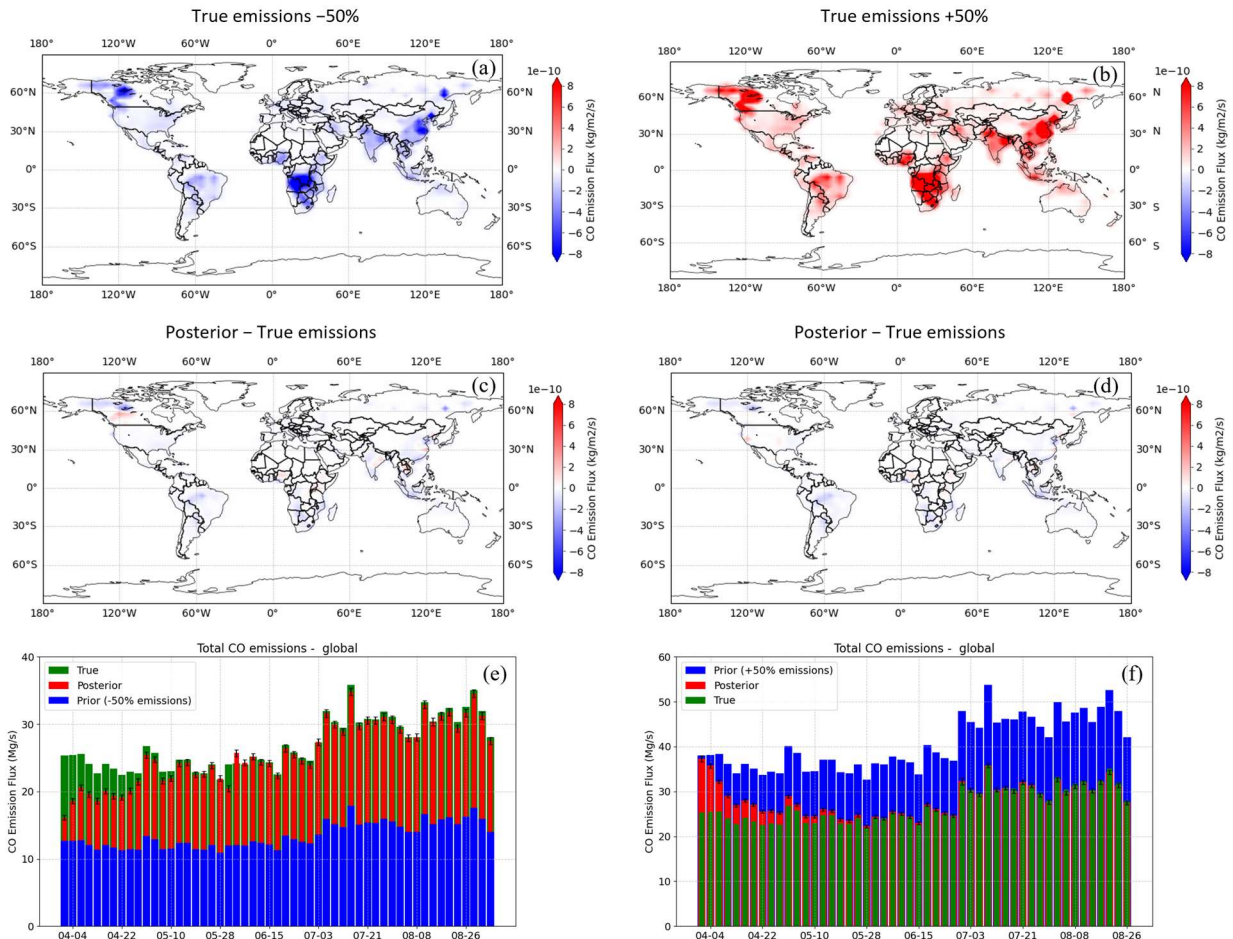


Figure A1: Two OSSEs within the twin experiments that start with the a priori of -50% CO emissions with respect to the true emissions as shown in (a) prior - true - prior and (c) posterior - true emissions; and with the a priori of $+50\%$ CO emissions with respect to the true emissions as shown in (b) prior - true - prior and (d) posterior - true emissions; time series of CO emissions in the a priori (blue), a posteriori (red), and true (green) for the OSSE with (e) -50% CO emissions and (f) $+50\%$ CO emissions.

Each OSSE setup involves multiple inversion runs. A “nature run” (without observation assimilation) is conducted to generate the “true” state of the concentration fields using GEOS-Chem with unperturbed emissions—the a priori emissions in the inversion with real observations. The true state is then mapped into the observation space by an observation operator, generating simulated observations that include added observation errors and under a perfect model transport assumption. A set of “control runs” assimilates these simulated observations. Each control run may vary in terms of perturbations in the magnitude of the emissions and/or the assimilation parameter range, depending on the experiment’s objectives. In our first OSSE, we use a combined set of simulated TROPOMI and TCCON observations and apply observation errors proportional to the retrieval errors, based on the ratio of simulated to retrieval XCO. Emissions in the control runs are perturbed by $\pm 50\%$ to evaluate the system's performance in recovering the true emissions.

Fig. A1 shows that the posteriors (Fig. A1 c,d) effectively capture and recover the spatial context of the CO emissions globally for both control runs, which have $\pm 50\%$ CO emissions in the prior (Fig. A1 a,b). Additionally, the time series of total CO emissions reveal that approximately 1.5 months (from the start of the assimilation) are required to constrain the magnitude and temporal variability of the global CO emissions (Fig. A1 e,f). In other experiments (not shown) where perturbations are applied only to major source regions (Fig. 2), we observe similar behaviour, even though the convergence rate for recovering true emissions slightly varies. For example, emissions are recovered relatively faster in North America (within ~ 3 weeks), followed by Europe (~ 4 weeks), likely due to a higher density of TCCON observations in these regions. This trend holds despite the nearly globally uniform spatiotemporal distribution of quality-filtered TROPOMI observations (i.e., super observations). In this experiment, we employ a set of previously optimized LETKF parameters and error statistics for background and observation errors, described below in Appendix B.

Appendix B: Comparing OSSEs with TROPOMI-only and TCCON-only inversions against joint TCCON and TROPOMI inversion

Similar OSSEs to those in Appendix A are conducted, using either only TROPOMI or only TCCON observations, with prior emissions perturbed by $\pm 50\%$ in one region at a time, as shown in Table B1. We compute the mean bias and standard deviation of OmF, along with the convergence time from the start of assimilation to recover the true emissions. Here, the forecast represents a model field, driven by either a priori or a posteriori emissions, mapped into the observation space. For the TROPOMI-only and TCCON-only inversions, we observe that biases and standard deviations generally increase across most cases; however, the inversion remains capable of recovering the true emissions. Compared to the joint TROPOMI and TCCON inversion, the convergence rate for the TROPOMI-only inversion slows down by up to a factor of two, depending on the perturbed region in the inversion. In contrast, convergence for the TCCON-only inversion can vary significantly. Specifically, the time to recover true emissions extends to 3 months in North America and 4 months in Europe, likely due to the lower spatial coverage of TCCON observations in these areas compared to TROPOMI. Additionally, we find that the TCCON-only inversion requires considerably more time to recover the true emissions globally and, especially, in Southern Hemisphere regions, such as Africa. The delay is primarily due to the limited number of TCCON sites in the Southern Hemisphere—only two sites are available for this study—and the extended time of inter-hemispheric exchange of air, which takes about 1 year (Jacob, 1999). As a result, sufficient information to constrain emissions in the Southern Hemisphere may not be achievable, especially within the limited 5-month period of inversion in this study.

Table B1: Mean bias and standard deviation of OmF for prior, posterior using TCCON-only, TROPOMI-only, and joint TROPOMI and TCCON (TROPOMI+TCCON) emissions. Each OSSE starts with -50% prior emissions in the specified region. Convergence times to recover the true emissions are shown from the start of assimilation.

Region	OSSE (-50% CO emissions)	OmF ^a mean bias (ppb)	OmF standard deviation (ppb)	Convergence time (month)
Global	Prior	-14.2	8.2	-
	TCCON-only	-2.5	4.7	$>6^b$
	TROPOMI-only	0.9	4.1	3.0
	TROPOMI+TCCON	-0.3	3.5	1.5
North America	Prior	-7.9	4.5	-
	TCCON-only	-0.5	2.6	3.0
	TROPOMI-only	0.7	2.8	2.0
	TROPOMI+TCCON	-0.2	2.1	1.0
Europe	Prior	-3.1	4.7	-
	TCCON-only	-0.4	2.1	4.0
	TROPOMI-only	-0.5	2.2	2.5
	TROPOMI+TCCON	-0.3	1.9	1.5
Africa	Prior	-6.2	5.5	-
	TCCON-only	-3.7	3.6	>6
	TROPOMI-only	0.6	2.3	2.0
	TROPOMI+TCCON	0.5	2.0	2.0

^a Forecast (F) in OmF uses a posteriori emissions for the assimilation run and a prior emissions for the prior (control) run.

^b It means that the true emissions is-are not fully recovered, $(|E_t^{\text{posterior}} - E_t^{\text{true}}|/E_t^{\text{true}}) \nless \delta$, within 6 months of inversion ($\delta = 2\%$).

To optimize the performance of the inversion, we employ OmF diagnostics to estimate key LETKF parameters. Specifically, we use the global mean bias and standard deviation of OmF over the full assimilation period to derive optimal values for parameters, such as the regularization factor γ , the inflation factor Δ , and the localization radius (r). We also configure essential setup elements, like the assimilation spin-up time, the burn-in duration, and the ensemble size to support efficient system operation. Our analysis yields the following optimal values: $\gamma_{TROPOMI} = 0.2$ and $\gamma_{TCCON} = 5$, $\Delta = 0.08$, $r = 500 \text{ km}$, with a minimum of three months for spin-up, one month for burn-in, and a minimum of 24 ensemble members—each chosen to minimize OmF statistics (e.g., see Fig. S4). Although we assume no transport modelling error, we apply an additional adjustment by inflating observation errors to offset this assumption. These configurations are consistently used across all OSSEs in this study.

Overall, our OSSEs results indicate that assimilating TCCON observations alone may not provide us with sufficient information to fully constrain the spatial context of CO emissions in regions in the Southern Hemisphere within the limited study period. However, in regions of the Northern Hemisphere, including North America and Europe, CO emissions are fully recovered within 2–3 months, likely due to the higher density of TCCON stations in these areas. In contrast, in the joint inversion, TROPOMI observations address the larger spatial biases, which typically exist in the model a priori, while TCCON measurements contribute finer constraints that enhance the representation of spatiotemporal variability.

Code and data availability. TROPOMI CO data can be downloaded from <https://doi.org/10.5270/S5P-bj3nry0> (Copernicus Sentinel-5P, 2024). The individual TCCON GGG2020 datasets used in this publication are cited in Table 1, and the references are included in the reference list. The TCCON data are available at <https://tccodata.org/2020> (last access: 6 July 2024) (Total Carbon Column Observing Network (TCCON) Team, 2022). The NDACC data are obtained as part of the Network for the Detection of Atmospheric Composition Change (NDACC) and are publicly available (see <https://www-air.larc.nasa.gov/missions/ndacc/data.html>; last access: 6 July, 2024). CO in situ measurements from WDCGG are available at <http://ds.data.jma.go.jp/gmd/wdcgg/> (last access: 1 July 2024). In situ aircraft CO measurements from Global Monitoring Laboratory of the National Oceanic and Atmospheric Administration (NOAA) are available at https://gml.noaa.gov/aftp/data/trace_gases/co/pfp/aircraft/ (last access: 1 July 2024). In situ tall tower measurements at ETL provided by Environment and Climate Change Canada are available at <https://gaw.kishou.go.jp/search/station#4007> (last access: 1 July 2024). GEOS-Chem version 14.1.1 source code is archived at <https://doi.org/10.5281/zenodo.7696651> (The International GEOS-Chem User Community, 2023), and MERRA-2 meteorology input data can be downloaded from WashU data portal at http://geoschemdata.wustl.edu/ExtData/GEOS_2x2.5/MERRA2/ (last access: 1 July 2024). The QFED emissions (version 2.5, release 1) data can be accessed from <http://geoschemdata.wustl.edu/ExtData/HEMCO/QFED/v2023-05/> (last access: 1 July 2024). The GBBEPx version 4 emissions data are available at <https://www.ospo.noaa.gov/pub/Blended/GBBEPx/> (last access: 1 July 2024). GFAS emissions (version 1.2) can be downloaded from <https://ads.atmosphere.copernicus.eu/datasets/cams-global-fire-emissions-gfas?tab=overview> (last access: 11 June 2024). The CFFEPS output is produced for ECCC's operational air quality forecast system (D. Kornic et al., 2024). The CFFEPS emissions code and the accompanying user manual are available at <https://zenodo.org/records/2579383> (last access: 1 July 2024) (Anderson and coauthors, 2019). The CHEEREIO source code is available at <https://github.com/drewpendergrass/CHEEREIO> (last access: 1 July 2024) (Pendergrass et al., 2024) and is documented at <https://cheereio.readthedocs.io> (last access: 22 August 2024). A forked repository of CHEEREIO used in this study, which ~~that~~ contains the TCCON CO and TROPOMI CO observation operators and the assimilation configuration, is available at <https://github.com/Sinavo/CHEEREIO> (last access: 22 August 2024).

Acknowledgements. We would like to thank the team behind the four BB emission inventories: QFED_v2.5r1, GBBEPx v4, GFASv1.2, and CFFEPSv4, for making their data available. We thank the World Data Centre for Greenhouse Gases

970 (WDCGG) for providing their CO data. We acknowledge all members of the Japan Meteorological Agency for operating the trace gas measuring systems at Minamitorishima, Yonagunijima, and Ryori stations. We thank the NOAA Greenhouse Gas Reference Network team and pilot Jared Chursinoff for their contributions to the aircraft data. The FTIR monitoring program at Jungfraujoch is primarily supported by the F.R.S. - FNRS (Brussels, Belgium) and the GAW-CH program of MeteoSwiss (Zürich, CH). The Paris site has received funding from Sorbonne Université, the French research center CNRS and the French
975 space agency CNES. Financial support from DGPAPA-UNAM (IN106024, G101225) was provided for the Mexican stations as well as the postdoctoral grant of Andrea Cadena-Caicedo. Tsukuba TCCON/NDACC and Rikubetsu TCCON sites are supported in part by the GOSAT series project. We thank the Canada Foundation for Innovation and NSERC for infrastructure and data analysis support for the TCCON station at ETL. Computations were performed on the Niagara supercomputer at the SciNet HPC Consortium. SciNet is funded by Innovation, Science and Economic Development Canada; the Digital Research
980 Alliance of Canada; the Ontario Research Fund: Research Excellence; and the University of Toronto.

Financial support. The Government of Canada's Environmental Damages Fund provided funding for this project under its Climate Action and Awareness Fund. Funding was also provided by the NOAA Global Monitoring Laboratory through Contract 1305M323PNRMJ0742.

985

Author contributions. The study was designed by SV, DBAJ, and DWu. SV performed the data analysis and prepared the original draft. DBAJ and DWu guided the work and edited the paper. DCP is the principal developer of CHEEREIO and SV contributed to the code development of TCCON and TROPOMI CO observation operators. JC, KA, and RSt contributed to the CFFEPS emissions. SK, and AZ contributed to GBBEPx emissions. DWu, POW, DFP, IMo, HO, NMD, FH, RSu, DWu, RK, OG, YT provided TCCON data. IMo, MVM, NJ, EM, AC, IMu contributed to NDACC measurements used. KM provided aircraft measurements. DWo and SR provided tall tower measurements at East Trout Lake. PC, CL, EK, TS, MS, RM contributed to WDCGG measurements. All authors reviewed and commented on the paper.

990

Competing interests. The authors declare that they have no conflict of interest

995

Disclaimer. Views expressed are those of authors and do not reflect those of NOAA or Department of Commerce.

References

1000 Abida, R., Attié, J.-L., El Amraoui, L., Ricaud, P., Lahoz, W., Eskes, H., Segers, A., Curier, L., de Haan, J., Kujanpää, J., Nijhuis, A. O., Tamminen, J., Timmermans, R., and Veeffkind, P.: Impact of spaceborne carbon monoxide observations from the S-5P platform on tropospheric composition analyses and forecasts, *Atmospheric Chem. Phys.*, 17, 1081–1103, <https://doi.org/10.5194/acp-17-1081-2017>, 2017.

- Alberti, C., Hase, F., Frey, M., Dubravica, D., Blumenstock, T., Dehn, A., Castracane, P., Surawicz, G., Harig, R., Baier, B. C., Bès, C., Bi, J., Boesch, H., Butz, A., Cai, Z., Chen, J., Crowell, S. M., Deutscher, N. M., Ene, D., Franklin, J. E., García, O., Griffith, D., Grouiez, B., Grutter, M., Hamdouni, A., Houweling, S., Humpage, N., Jacobs, N., Jeong, S., Joly, L., Jones, N. B., Jouglet, D., Kivi, R., Kleinschek, R., Lopez, M., Medeiros, D. J., Morino, I., Mostafavipak, N., Müller, A., Ohyama, H., Palmer, P. I., Pathakoti, M., Pollard, D. F., Raffalski, U., Ramonet, M., Ramsay, R., Sha, M. K., Shiomi, K., Simpson, W., Stremme, W., Sun, Y., Tanimoto, H., Té, Y., Tsidu, G. M., Velazco, V. A., Vogel, F., Watanabe, M., Wei, C., Wunch, D., Yamasoe, M., Zhang, L., and Orphal, J.: Improved calibration procedures for the EM27/SUN spectrometers of the COLlaborative Carbon Column Observing Network (COCCON), *Atmospheric Meas. Tech.*, 15, 2433–2463, <https://doi.org/10.5194/amt-15-2433-2022>, 2022.
- Anderson, K. and cast of thousands: Canadian Fire Emissions Prediction System (CFFEPS), , <https://doi.org/10.5281/zenodo.2579383>, 2019.
- Anderson, K. R., Pankratz, A., and Mooney, C.: A thermodynamic approach to estimating smoke plume heights., 2011.
- Andreae, M. O.: Emission of trace gases and aerosols from biomass burning – an updated assessment, *Atmospheric Chem. Phys.*, 19, 8523–8546, <https://doi.org/10.5194/acp-19-8523-2019>, 2019.
- Apituley, A.: Sentinel-5 precursor/TROPOMI Level 2 Product User Manual Carbon Monoxide, n.d.
- Arellano Jr., A. F., Kasibhatla, P. S., Giglio, L., van der Werf, G. R., Randerson, J. T., and Collatz, G. J.: Time-dependent inversion estimates of global biomass-burning CO emissions using Measurement of Pollution in the Troposphere (MOPITT) measurements, *J. Geophys. Res. Atmospheres*, 111, <https://doi.org/10.1029/2005JD006613>, 2006.
- Aschi, M. and Largo, A.: Reactivity of gaseous protonated ozone: a computational investigation on the carbon monoxide oxidation reaction, *Int. J. Mass Spectrom.*, 228, 613–627, [https://doi.org/10.1016/S1387-3806\(03\)00134-9](https://doi.org/10.1016/S1387-3806(03)00134-9), 2003.
- Bègue, N., Baron, A., Krysztofiak, G., Berthet, G., Kloss, C., Jégou, F., Khaykin, S., Ranaivombola, M., Millet, T., Portafaix, T., Dufлот, V., Keckhut, P., Vèremes, H., Payen, G., Sha, M. K., Coheur, P.-F., Clerbaux, C., Sicard, M., Sakai, T., Querel, R., Liley, B., Smale, D., Morino, I., Uchino, O., Nagai, T., Smale, P., Robinson, J., and Bencherif, H.: Evidence of a dual African and Australian biomass burning influence on the vertical distribution of aerosol and carbon monoxide over the southwest Indian Ocean basin in early 2020, *Atmospheric Chem. Phys.*, 24, 8031–8048, <https://doi.org/10.5194/acp-24-8031-2024>, 2024.
- Bey, I., Jacob, D. J., Yantosca, R. M., Logan, J. A., Field, B. D., Fiore, A. M., Li, Q., Liu, H. Y., Mickley, L. J., and Schultz, M. G.: Global modeling of tropospheric chemistry with assimilated meteorology: Model description and evaluation, *J. Geophys. Res. Atmospheres*, 106, 23073–23095, <https://doi.org/10.1029/2001JD000807>, 2001.
- Bisht, J. S. H., Patra, P. K., Takigawa, M., Sekiya, T., Kanaya, Y., Saitoh, N., and Miyazaki, K.: Estimation of CH₄ emission based on an advanced 4D-LETKF assimilation system, *Geosci. Model Dev.*, 16, 1823–1838, <https://doi.org/10.5194/gmd-16-1823-2023>, 2023.
- Bocquet, M., Elbern, H., Eskes, H., Hirtl, M., Žabkar, R., Carmichael, G. R., Flemming, J., Inness, A., Pagowski, M., Pérez Camacho, J. L., Saide, P. E., San Jose, R., Sofiev, M., Vira, J., Baklanov, A., Carnevale, C., Grell, G., and Seigneur, C.: Data assimilation in atmospheric chemistry models: current status and future prospects for coupled chemistry meteorology models, *Atmospheric Chem. Phys.*, 15, 5325–5358, <https://doi.org/10.5194/acp-15-5325-2015>, 2015.
- Borsdorff, T., Aan de Brugh, J., Hu, H., Aben, I., Hasekamp, O., and Landgraf, J.: Measuring Carbon Monoxide With TROPOMI: First Results and a Comparison With ECMWF-IFS Analysis Data, *Geophys. Res. Lett.*, 45, 2826–2832, <https://doi.org/10.1002/2018GL077045>, 2018.

- 1040 Borsdorff, T., aan de Brugh, J., Schneider, A., Lorente, A., Birk, M., Wagner, G., Kivi, R., Hase, F., Feist, D. G., Sussmann, R., Rettinger, M., Wunch, D., Warneke, T., and Landgraf, J.: Improving the TROPOMI CO data product: update of the spectroscopic database and destripping of single orbits, *Atmospheric Meas. Tech.*, 12, 5443–5455, <https://doi.org/10.5194/amt-12-5443-2019>, 2019.
- 1045 Borsdorff, T., García Reynoso, A., Maldonado, G., Mar-Morales, B., Stremme, W., Grutter, M., and Landgraf, J.: Monitoring CO emissions of the metropolis Mexico City using TROPOMI CO observations, *Atmospheric Chem. Phys.*, 20, 15761–15774, <https://doi.org/10.5194/acp-20-15761-2020>, 2020.
- Borsdorff, T., Campos, T., Kille, N., Zarzana, K. J., Volkamer, R., and Landgraf, J.: Vertical information of CO from TROPOMI total column measurements in context of the CAMS-IFS data assimilation scheme, *Atmospheric Meas. Tech.*, 16, 3027–3038, <https://doi.org/10.5194/amt-16-3027-2023>, 2023.
- 1050 Bukosa, B., Fisher, J. A., Deutscher, N. M., and Jones, D. B. A.: A Coupled CH₄, CO and CO₂ Simulation for Improved Chemical Source Modeling, *Atmosphere*, 14, 764, <https://doi.org/10.3390/atmos14050764>, 2023.
- 1055 Byrne, B., Liu, J., Lee, M., Baker, I., Bowman, K. W., Deutscher, N. M., Feist, D. G., Griffith, D. W. T., Iraci, L. T., Kiel, M., Kimball, J. S., Miller, C. E., Morino, I., Parazoo, N. C., Petri, C., Roehl, C. M., Sha, M. K., Strong, K., Velazco, V. A., Wennberg, P. O., and Wunch, D.: Improved Constraints on Northern Extratropical CO₂ Fluxes Obtained by Combining Surface-Based and Space-Based Atmospheric CO₂ Measurements, *J. Geophys. Res. Atmospheres*, 125, e2019JD032029, <https://doi.org/10.1029/2019JD032029>, 2020.
- Byrne, B., Liu, J., Bowman, K. W., Pascolini-Campbell, M., Chatterjee, A., Pandey, S., Miyazaki, K., van der Werf, G. R., Wunch, D., Wennberg, P. O., Roehl, C. M., and Sinha, S.: Carbon emissions from the 2023 Canadian wildfires, *Nature*, 633, 835–839, <https://doi.org/10.1038/s41586-024-07878-z>, 2024.
- 1060 Cascio, W. E.: Wildland fire smoke and human health, *Sci. Total Environ.*, 624, 586–595, <https://doi.org/10.1016/j.scitotenv.2017.12.086>, 2018.
- Chen, B., Zhang, H., Coops, N. C., Fu, D., Worthy, D. E. J., Xu, G., and Black, T. A.: Assessing scalar concentration footprint climatology and land surface impacts on tall-tower CO₂ concentration measurements in the boreal forest of central Saskatchewan, Canada, *Theor. Appl. Climatol.*, 118, 115–132, <https://doi.org/10.1007/s00704-013-1038-2>, 2014.
- 1065 Chen, J., Anderson, K., Pavlovic, R., Moran, M. D., Englefield, P., Thompson, D. K., Munoz-Alpizar, R., and Landry, H.: The FireWork v2.0 air quality forecast system with biomass burning emissions from the Canadian Forest Fire Emissions Prediction System v2.03, *Geosci. Model Dev.*, 12, 3283–3310, <https://doi.org/10.5194/gmd-12-3283-2019>, 2019.
- 1070 Chen, J., Yao, Q., Chen, Z., Li, M., Hao, Z., Liu, C., Zheng, W., Xu, M., Chen, X., Yang, J., Lv, Q., and Gao, B.: The Fengyun-3D (FY-3D) global active fire product: principle, methodology and validation, *Earth Syst. Sci. Data*, 14, 3489–3508, <https://doi.org/10.5194/essd-14-3489-2022>, 2022.
- Chen, Y., Morton, D. C., Andela, N., van der Werf, G. R., Giglio, L., and Randerson, J. T.: A pan-tropical cascade of fire driven by El Niño/Southern Oscillation, *Nat. Clim. Change*, 7, 906–911, <https://doi.org/10.1038/s41558-017-0014-8>, 2017.
- 1075 Chen, Z., Jacob, D. J., Gautam, R., Omara, M., Stavins, R. N., Stowe, R. C., Nesser, H., Sulprizio, M. P., Lorente, A., Varon, D. J., Lu, X., Shen, L., Qu, Z., Pendergrass, D. C., and Hancock, S.: Satellite quantification of methane emissions and oil–gas methane intensities from individual countries in the Middle East and North Africa: implications for climate action, *Atmospheric Chem. Phys.*, 23, 5945–5967, <https://doi.org/10.5194/acp-23-5945-2023>, 2023.

- Chevallier, F., Deutscher, N. M., Conway, T. J., Ciais, P., Ciattaglia, L., Dohe, S., Fröhlich, M., Gomez-Pelaez, A. J., Griffith, D., Hase, F., Haszpra, L., Krummel, P., Kyrö, E., Labuschagne, C., Langenfelds, R., Machida, T., Maignan, F., Matsueda, H., Morino, I., Notholt, J., Ramonet, M., Sawa, Y., Schmidt, M., Sherlock, V., Steele, P., Strong, K., Sussmann, R., Wennberg, P., Wofsy, S., Worth, D., Wunch, D., and Zimnoch, M.: Global CO₂ fluxes inferred from surface air-sample measurements and from TCCON retrievals of the CO₂ total column, *Geophys. Res. Lett.*, 38, <https://doi.org/10.1029/2011GL049899>, 2011.
- Copernicus Sentinel-5P: (processed by ESA), TROPOMI Level 2 Carbon Monoxide total column products (02), <https://doi.org/10.5270/S5P-bj3nry0>, 2024.
- Cristofanelli, P., Busetto, M., Calzolari, F., Ammoscato, I., Gulli, D., Dinoi, A., Calidonna, C. R., Contini, D., Sferlazzo, D., Di Iorio, T., Piacentino, S., Marinoni, A., Maione, M., and Bonasoni, P.: Investigation of reactive gases and methane variability in the coastal boundary layer of the central Mediterranean basin, *Elem. Sci. Anthr.*, 5, 12, <https://doi.org/10.1525/elementa.216>, 2017.
- Cristofanelli, P., Trisolino, P., Calzolari, F., Busetto, M., Calidonna, C. R., Amendola, S., Arduini, J., Fratticioli, C., Hundal, R. A., Maione, M., Marcucci, F., Marinoni, A., Montaguti, S., Renzi, L., Roccato, F., Bonasoni, P., and Putero, D.: Influence of wildfire emissions to carbon dioxide (CO₂) observed at the Mt. Cimone station (Italy, 2165 m asl): A multi-year investigation, *Atmos. Environ.*, 330, 120577, <https://doi.org/10.1016/j.atmosenv.2024.120577>, 2024.
- Csiszar, I., Schroeder, W., Giglio, L., Mikles, V., and Tsidulko, M.: TITLE: NDE ACTIVE FIRE EDR EXTERNAL USERS MANUAL VERSION 1.6, n.d.
- D. Kornic, S. Ménard, C. Stroud, C. Zaganescu, R. Munoz-Alpizar, R. Ménard, S. Antonopoulos, and J. Montpetit: Regional Air Quality Deterministic Prediction System (RAQDPS), 2024.
- De Mazière, M., Thompson, A. M., Kurylo, M. J., Wild, J. D., Bernhard, G., Blumenstock, T., Braathen, G. O., Hannigan, J. W., Lambert, J.-C., Leblanc, T., McGee, T. J., Nedoluha, G., Petropavlovskikh, I., Seckmeyer, G., Simon, P. C., Steinbrecht, W., and Strahan, S. E.: The Network for the Detection of Atmospheric Composition Change (NDACC): history, status and perspectives, *Atmospheric Chem. Phys.*, 18, 4935–4964, <https://doi.org/10.5194/acp-18-4935-2018>, 2018.
- Deeter, M. N., Emmons, L. K., Francis, G. L., Edwards, D. P., Gille, J. C., Warner, J. X., Khattatov, B., Ziskin, D., Lamarque, J.-F., Ho, S.-P., Yudin, V., Attié, J.-L., Packman, D., Chen, J., Mao, D., and Drummond, J. R.: Operational carbon monoxide retrieval algorithm and selected results for the MOPITT instrument, *J. Geophys. Res. Atmospheres*, 108, <https://doi.org/10.1029/2002JD003186>, 2003.
- Deeter, M. N., Edwards, D. P., Gille, J. C., and Worden, H. M.: Information content of MOPITT CO profile retrievals: Temporal and geographical variability, *J. Geophys. Res. Atmospheres*, 120, 12723–12738, <https://doi.org/10.1002/2015JD024024>, 2015.
- Deutscher, N. M., Griffith, D. W. T., Paton-Walsh, C., Jones, N. B., Velazco, V. A., Wilson, S. R., Macatangay, R. C., Kettlewell, G. C., Buchholz, R. R., Riggensbach, M. O., Bukosa, B., John, S. S., Walker, B. T., and Nawaz, H.: TCCON data from Wollongong (AU), Release GGG2020.R0 (R0), <https://doi.org/10.14291/tcon.ggg2020.wollongong01.R0>, 2023.
- Di Giuseppe, F., Benedetti, A., Coughlan, R., Vitolo, C., and Vuckovic, M.: A Global Bottom-Up Approach to Estimate Fuel Consumed by Fires Using Above Ground Biomass Observations, *Geophys. Res. Lett.*, 48, e2021GL095452, <https://doi.org/10.1029/2021GL095452>, 2021.
- Edwards, D. P., Pétron, G., Novelli, P. C., Emmons, L. K., Gille, J. C., and Drummond, J. R.: Southern Hemisphere carbon monoxide interannual variability observed by Terra/Measurement of Pollution in the Troposphere (MOPITT), *J. Geophys. Res. Atmospheres*, 111, <https://doi.org/10.1029/2006JD007079>, 2006.

- Eko Cahyono, W., Parikesit, Joy, B., Setyawati, W., and Mahdi, R.: Projection of CO₂ emissions in Indonesia, *Mater. Today Proc.*, 63, S438–S444, <https://doi.org/10.1016/j.matpr.2022.04.091>, 2022.
- Eskes, H. J., Velthoven, P. F. J. V., Valks, P. J. M., and Kelder, H. M.: Assimilation of GOME total-ozone satellite observations in a three-dimensional tracer-transport model, *Q. J. R. Meteorol. Soc.*, 129, 1663–1681, <https://doi.org/10.1256/qj.02.14>, 2003.
- 1120 Feng, L., Palmer, P. I., Bösch, H., and Dance, S.: Estimating surface CO₂ fluxes from space-borne CO₂ dry air mole fraction observations using an ensemble Kalman Filter, *Atmospheric Chem. Phys.*, 9, 2619–2633, <https://doi.org/10.5194/acp-9-2619-2009>, 2009.
- Feng, S., Jiang, F., Wu, Z., Wang, H., Ju, W., and Wang, H.: CO Emissions Inferred From Surface CO Observations Over China in December 2013 and 2017, *J. Geophys. Res. Atmospheres*, 125, e2019JD031808, <https://doi.org/10.1029/2019JD031808>, 2020.
- 1125 Filizzola, C., Falconieri, A., Lacava, T., Marchese, F., Masiello, G., Mazzeo, G., Pergola, N., Pietrapertosa, C., Serio, C., and Tramutoli, V.: Fire Characterization by Using an Original RST-Based Approach for Fire Radiative Power (FRP) Computation, *Fire*, 6, 48, <https://doi.org/10.3390/fire6020048>, 2023.
- Fischer, E. V., Jacob, D. J., Yantosca, R. M., Sulprizio, M. P., Millet, D. B., Mao, J., Paulot, F., Singh, H. B., Roiger, A., Ries, L., Talbot, R. W., Dzepina, K., and Pandey Deolal, S.: Atmospheric peroxyacetyl nitrate (PAN): a global budget and source attribution, *Atmospheric Chem. Phys.*, 14, 2679–2698, <https://doi.org/10.5194/acp-14-2679-2014>, 2014.
- 1130 Fisher, J. A., Murray, L. T., Jones, D. B. A., and Deutscher, N. M.: Improved method for linear carbon monoxide simulation and source attribution in atmospheric chemistry models illustrated using GEOS-Chem v9, *Geosci. Model Dev.*, 10, 4129–4144, <https://doi.org/10.5194/gmd-10-4129-2017>, 2017.
- 1135 Fowler, D., Amann, M., Anderson, R., Ashmore, M., Cox, P., Depledge, M., Derwent, D., Grennfelt, P., Hewitt, N., Hov, O., Jenkin, M., Kelly, F., Liss, P. S., Pilling, M., Pyle, J., Slingo, J., and Stevenson, D.: Ground-level ozone in the 21st century: Future trends, impacts and policy implications, 2008.
- Frey, M., Sha, M. K., Hase, F., Kiel, M., Blumenstock, T., Harig, R., Surawicz, G., Deutscher, N. M., Shiomi, K., Franklin, J. E., Bösch, H., Chen, J., Grutter, M., Ohyama, H., Sun, Y., Butz, A., Mengistu Tsidu, G., Ene, D., Wunch, D., Cao, Z., Garcia, O., Ramonet, M., Vogel, F., and Orphal, J.: Building the Collaborative Carbon Column Observing Network (COCCON): long-term stability and ensemble performance of the EM27/SUN Fourier transform spectrometer, *Atmospheric Meas. Tech.*, 12, 1513–1530, <https://doi.org/10.5194/amt-12-1513-2019>, 2019.
- 1140 García, O. E., Schneider, M., Herkommer, B., Gross, J., Hase, F., Blumenstock, T., and Sepúlveda, E.: TCCON data from Izana (ES), Release GGG2020.R1 (R1), <https://doi.org/10.14291/tccon.ggg2020.izana01.R1>, 2022.
- 1145 Gaubert, B., Edwards, D. P., Anderson, J. L., Arellano, A. F., Barré, J., Buchholz, R. R., Darras, S., Emmons, L. K., Fillmore, D., Granier, C., Hannigan, J. W., Ortega, I., Raeder, K., Soulié, A., Tang, W., Worden, H. M., and Ziskin, D.: Global Scale Inversions from MOPITT CO and MODIS AOD, *Remote Sens.*, 15, 4813, <https://doi.org/10.3390/rs15194813>, 2023.
- Gelaro, R., McCarty, W., Suárez, M. J., Todling, R., Molod, A., Takacs, L., Randles, C. A., Darmenov, A., Bosilovich, M. G., Reichle, R., Wargan, K., Coy, L., Cullather, R., Draper, C., Akella, S., Buchard, V., Conaty, A., Silva, A. M. da, Gu, W., Kim, G.-K., Koster, R., Lucchesi, R., Merkova, D., Nielsen, J. E., Partyka, G., Pawson, S., Putman, W., Rienecker, M., Schubert, S. D., Sienkiewicz, M., and Zhao, B.: The Modern-Era Retrospective Analysis for Research and Applications, Version 2 (MERRA-2), *J. Clim.*, 30, 5419–5454, <https://doi.org/10.1175/JCLI-D-16-0758.1>, 2017.
- 1150

- Ghil, M. and Mo, K.: Intraseasonal Oscillations in the Global Atmosphere. Part I: Northern Hemisphere and Tropics, *J. Atmospheric Sci.*, 48, 752–779, [https://doi.org/10.1175/1520-0469\(1991\)048%253C0752:IOITGA%253E2.0.CO;2](https://doi.org/10.1175/1520-0469(1991)048%253C0752:IOITGA%253E2.0.CO;2), 1991.
- 1155 Goudar, M., Anema, J. C. S., Kumar, R., Borsdorff, T., and Landgraf, J.: Plume detection and emission estimate for biomass burning plumes from TROPOMI carbon monoxide observations using APE v1.1, *Geosci. Model Dev.*, 16, 4835–4852, <https://doi.org/10.5194/gmd-16-4835-2023>, 2023.
- 1160 Griffin, D., Sioris, C., Chen, J., Dickson, N., Kovachik, A., de Graaf, M., Nanda, S., Veefkind, P., Dammers, E., McLinden, C. A., Makar, P., and Akingunola, A.: The 2018 fire season in North America as seen by TROPOMI: aerosol layer height intercomparisons and evaluation of model-derived plume heights, *Atmospheric Meas. Tech.*, 13, 1427–1445, <https://doi.org/10.5194/amt-13-1427-2020>, 2020.
- Griffin, D., Chen, J., Anderson, K., Makar, P., McLinden, C. A., Dammers, E., and Fogal, A.: Biomass burning CO emissions: exploring insights through TROPOMI-derived emissions and emission coefficients, *Atmospheric Chem. Phys.*, 24, 10159–10186, <https://doi.org/10.5194/acp-24-10159-2024>, 2024.
- 1165 Grutter, M., Basaldud, R., Rivera, C., Harig, R., Junkerman, W., Caetano, E., and Delgado-Granados, H.: SO₂ emissions from Popocatepetl volcano: emission rates and plume imaging using optical remote sensing techniques, *Atmospheric Chem. Phys.*, 8, 6655–6663, <https://doi.org/10.5194/acp-8-6655-2008>, 2008.
- 1170 Guenther, A. B., Jiang, X., Heald, C. L., Sakulyanontvittaya, T., Duhl, T., Emmons, L. K., and Wang, X.: The Model of Emissions of Gases and Aerosols from Nature version 2.1 (MEGAN2.1): an extended and updated framework for modeling biogenic emissions, *Geosci. Model Dev.*, 5, 1471–1492, <https://doi.org/10.5194/gmd-5-1471-2012>, 2012.
- Hakami, A., Henze, D. K., Seinfeld, J. H., Chai, T., Tang, Y., Carmichael, G. R., and Sandu, A.: Adjoint inverse modeling of black carbon during the Asian Pacific Regional Aerosol Characterization Experiment, *J. Geophys. Res. Atmospheres*, 110, <https://doi.org/10.1029/2004JD005671>, 2005.
- 1175 Hase, F., Herkommer, B., Groß, J., Blumenstock, T., Kiel, M. ä., and Dohe, S.: TCCON data from Karlsruhe (DE), Release GGG2020.R1 (R1), <https://doi.org/10.14291/tccon.ggg2020.karlsruhe01.R1>, 2023.
- Hasekamp, O., Lorente, A., Hu, H., and Butz, A.: Algorithm Theoretical Baseline Document for Sentinel-5 Precursor Methane Retrieval, 2022.
- 1180 Heald, C. L., Jacob, D. J., Jones, D. B. A., Palmer, P. I., Logan, J. A., Streets, D. G., Sachse, G. W., Gille, J. C., Hoffman, R. N., and Nehr Korn, T.: Comparative inverse analysis of satellite (MOPITT) and aircraft (TRACE-P) observations to estimate Asian sources of carbon monoxide, *J. Geophys. Res. Atmospheres*, 109, <https://doi.org/10.1029/2004JD005185>, 2004.
- Hedelius, J. K., Toon, G. C., Buchholz, R. R., Iraci, L. T., Podolske, J. R., Roehl, C. M., Wennberg, P. O., Worden, H. M., and Wunch, D.: Regional and Urban Column CO Trends and Anomalies as Observed by MOPITT Over 16 Years, *J. Geophys. Res. Atmospheres*, 126, e2020JD033967, <https://doi.org/10.1029/2020JD033967>, 2021.
- 1185 Hoesly, R. M., Smith, S. J., Feng, L., Klimont, Z., Janssens-Maenhout, G., Pitkanen, T., Seibert, J. J., Vu, L., Andres, R. J., Bolt, R. M., Bond, T. C., Dawidowski, L., Kholod, N., Kurokawa, J., Li, M., Liu, L., Lu, Z., Moura, M. C. P., O'Rourke, P. R., and Zhang, Q.: Historical (1750–2014) anthropogenic emissions of reactive gases and aerosols from the Community Emissions Data System (CEDS), *Geosci. Model Dev.*, 11, 369–408, <https://doi.org/10.5194/gmd-11-369-2018>, 2018.
- Holloway, T., Levy, H., and Kasibhatla, P.: Global distribution of carbon monoxide, *J. Geophys. Res. Atmospheres*, 105, 12123–12147, <https://doi.org/10.1029/1999JD901173>, 2000.

- 1190 Hooghiemstra, P. B., Krol, M. C., Meirink, J. F., Bergamaschi, P., van der Werf, G. R., Novelli, P. C., Aben, I., and Röckmann, T.: Optimizing global CO emission estimates using a four-dimensional variational data assimilation system and surface network observations, *Atmospheric Chem. Phys.*, 11, 4705–4723, <https://doi.org/10.5194/acp-11-4705-2011>, 2011.
- 1195 Hueglin, C., Buchmann, B., Steinbacher, M., and Emmenegger, L.: The Swiss National Air Pollution Monitoring Network (NABEL) – Bridging Science and Environmental Policy, *CHIMIA*, 78, 722–727, <https://doi.org/10.2533/chimia.2024.722>, 2024.
- Hundal, R. A., Annadate, S., Cesari, R., Collalti, A., Maione, M., and Cristofanelli, P.: Emissions of climate-altering species from open vegetation fires in the Mediterranean region - A review on methods and data, *Sci. Total Environ.*, 957, 177713, <https://doi.org/10.1016/j.scitotenv.2024.177713>, 2024.
- 1200 Hunt, B. R., Kostelich, E. J., and Szunyogh, I.: Efficient data assimilation for spatiotemporal chaos: A local ensemble transform Kalman filter, *Phys. Nonlinear Phenom.*, 230, 112–126, <https://doi.org/10.1016/j.physd.2006.11.008>, 2007.
- Inness, A., Aben, I., Ades, M., Borsdorff, T., Flemming, J., Jones, L., Landgraf, J., Langerock, B., Nedelec, P., Parrington, M., and Ribas, R.: Assimilation of S5P/TROPOMI carbon monoxide data with the global CAMS near-real-time system, *Atmospheric Chem. Phys.*, 22, 14355–14376, <https://doi.org/10.5194/acp-22-14355-2022>, 2022.
- 1205 Iraci, L. T., Podolske, J. R., Roehl, C., Wennberg, P. O., Blavier, J.-F., Allen, N., Wunch, D., and Osterman, G. B.: TCCON data from Edwards (US), Release GGG2020.R0 (R0), <https://doi.org/10.14291/tcon.ggg2020.edwards01.R0>, 2022.
- Jacob, D. J.: *Introduction to Atmospheric Chemistry*, Princeton University Press, 1999.
- Jain, P., Barber, Q. E., Taylor, S. W., Whitman, E., Castellanos Acuna, D., Boulanger, Y., Chavardès, R. D., Chen, J., Englefield, P., Flannigan, M., Girardin, M. P., Hanes, C. C., Little, J., Morrison, K., Skakun, R. S., Thompson, D. K., Wang, X., and Parisien, M.-A.: Drivers and Impacts of the Record-Breaking 2023 Wildfire Season in Canada, *Nat. Commun.*, 15, 6764, <https://doi.org/10.1038/s41467-024-51154-7>, 2024.
- 1210 Jegasothy, E., Hanigan, I. C., Van Buskirk, J., Morgan, G. G., Jalaludin, B., Johnston, F. H., Guo, Y., and Broome, R. A.: Acute health effects of bushfire smoke on mortality in Sydney, Australia, *Environ. Int.*, 171, 107684, <https://doi.org/10.1016/j.envint.2022.107684>, 2023.
- 1215 Jiang, Z., Worden, J. R., Worden, H., Deeter, M., Jones, D. B. A., Arellano, A. F., and Henze, D. K.: A 15-year record of CO emissions constrained by MOPITT CO observations, *Atmospheric Chem. Phys.*, 17, 4565–4583, <https://doi.org/10.5194/acp-17-4565-2017>, 2017.
- 1220 Jin, Y., Liu, Y., Lu, X., Chen, X., Shen, A., Wang, H., Cui, Y., Xu, Y., Li, S., Liu, J., Zhang, M., Ma, Y., and Fan, Q.: Measurement report: Assessing the impacts of emission uncertainty on aerosol optical properties and radiative forcing from biomass burning in peninsular Southeast Asia, *Atmospheric Chem. Phys.*, 24, 367–395, <https://doi.org/10.5194/acp-24-367-2024>, 2024.
- Jones, D. B. A., Bowman, K. W., Palmer, P. I., Worden, J. R., Jacob, D. J., Hoffman, R. N., Bey, I., and Yantosca, R. M.: Potential of observations from the Tropospheric Emission Spectrometer to constrain continental sources of carbon monoxide, *J. Geophys. Res. Atmospheres*, 108, <https://doi.org/10.1029/2003JD003702>, 2003.
- 1225 Jones, D. B. A., Bowman, K. W., Logan, J. A., Heald, C. L., Liu, J., Luo, M., Worden, J., and Drummond, J.: The zonal structure of tropical O₃ and CO as observed by the Tropospheric Emission Spectrometer in November 2004 – Part 1: Inverse modeling of CO emissions, *Atmospheric Chem. Phys.*, 9, 3547–3562, <https://doi.org/10.5194/acp-9-3547-2009>, 2009a.

- 1230 Jones, M. W., Kelley, D. I., Burton, C. A., Di Giuseppe, F., Barbosa, M. L. F., Brambleby, E., Hartley, A. J., Lombardi, A., Mataveli, G., McNorton, J. R., Spuler, F. R., Wessel, J. B., Abatzoglou, J. T., Anderson, L. O., Andela, N., Archibald, S., Armenteras, D., Burke, E., Carmenta, R., Chuvieco, E., Clarke, H., Doerr, S. H., Fernandes, P. M., Giglio, L., Hamilton, D. S., Hantson, S., Harris, S., Jain, P., Kolden, C. A., Kurvits, T., Lampe, S., Meier, S., New, S., Parrington, M., Perron, M. M. G., Qu, Y., Ribeiro, N. S., Saharjo, B. H., San-Miguel-Ayanz, J., Shuman, J. K., Tanpipat, V., van der Werf, G. R., Veraverbeke, S., and Xanthopoulos, G.: State of Wildfires 2023–2024, *Earth Syst. Sci. Data*, 16, 3601–3685, <https://doi.org/10.5194/essd-16-3601-2024>, 2024.
- 1235 Jones, N. B., Riedel, K., Allan, W., Wood, S., Palmer, P. I., Chance, K., and Notholt, J.: Long-term tropospheric formaldehyde concentrations deduced from ground-based fourier transform solar infrared measurements, *Atmospheric Chem. Phys.*, 9, 7131–7142, <https://doi.org/10.5194/acp-9-7131-2009>, 2009b.
- Kaiser, J. W., Heil, A., Andreae, M. O., Benedetti, A., Chubarova, N., Jones, L., Morcrette, J.-J., Razinger, M., Schultz, M. G., Suttie, M., and van der Werf, G. R.: Biomass burning emissions estimated with a global fire assimilation system based on observed fire radiative power, *Biogeosciences*, 9, 527–554, <https://doi.org/10.5194/bg-9-527-2012>, 2012.
- 1240 Kasibhatla, P., Arellano, A., Logan, J. A., Palmer, P. I., and Novelli, P.: Top-down estimate of a large source of atmospheric carbon monoxide associated with fuel combustion in Asia, *Geophys. Res. Lett.*, 29, 6-1-6-4, <https://doi.org/10.1029/2002GL015581>, 2002.
- 1245 Kerzenmacher, T., Dils, B., Kumps, N., Blumenstock, T., Clerbaux, C., Coheur, P.-F., Demoulin, P., García, O., George, M., Griffith, D. W. T., Hase, F., Hadji-Lazaro, J., Hurtmans, D., Jones, N., Mahieu, E., Notholt, J., Paton-Walsh, C., Raffalski, U., Ridder, T., Schneider, M., Servais, C., and De Mazière, M.: Validation of IASI FORLI carbon monoxide retrievals using FTIR data from NDACC, *Atmospheric Meas. Tech.*, 5, 2751–2761, <https://doi.org/10.5194/amt-5-2751-2012>, 2012.
- 1250 Kim, E., Kim, B.-U., Kim, H. C., Liu, Y., Kang, Y. H., Jacob, D. J., Kim, Y. P., Woo, J.-H., Kim, J., Wang, S., Yoo, C., Bae, C., Kim, Y., and Kim, S.: North Korean CO emissions reconstruction using DMZ ground observations, TROPOMI space-borne data, and the CMAQ air quality model, *Sci. Total Environ.*, 921, 171059, <https://doi.org/10.1016/j.scitotenv.2024.171059>, 2024.
- Kirago, L., Gustafsson, Ö., Gaita, S. M., Haslett, S. L., Gatari, M. J., Popa, M. E., Röckmann, T., Zellweger, C., Steinbacher, M., Klausen, J., Félix, C., Njiru, D., and Andersson, A.: Sources and long-term variability of carbon monoxide at Mount Kenya and in Nairobi, *Atmospheric Chem. Phys.*, 23, 14349–14357, <https://doi.org/10.5194/acp-23-14349-2023>, 2023.
- 1255 Kivi, R., Heikkinen, P., and Kyrö, E.: TCCON data from Sodankylä (FI), Release GGG2020.R0 (R0), <https://doi.org/10.14291/tccon.ggg2020.sodankyla01.R0>, 2022.
- Kolden, C. A., Abatzoglou, J. T., Jones, M. W., and Jain, P.: Wildfires in 2023, *Nat. Rev. Earth Environ.*, 5, 238–240, <https://doi.org/10.1038/s43017-024-00544-y>, 2024.
- 1260 Kopacz, M., Jacob, D. J., Fisher, J. A., Logan, J. A., Zhang, L., Megretskaia, I. A., Yantosca, R. M., Singh, K., Henze, D. K., Burrows, J. P., Buchwitz, M., Khlystova, I., McMillan, W. W., Gille, J. C., Edwards, D. P., Eldering, A., Thouret, V., and Nedelec, P.: Global estimates of CO sources with high resolution by adjoint inversion of multiple satellite datasets (MOPITT, AIRS, SCIAMACHY, TES), *Atmospheric Chem. Phys.*, 10, 855–876, <https://doi.org/10.5194/acp-10-855-2010>, 2010.
- Koster, R. D., Darmenov, A. S., and da Silva, A. M.: The Quick Fire Emissions Dataset (QFED): Documentation of Versions 2.1, 2.2 and 2.4: Technical Report Series on Global Modeling and Data Assimilation - Volume 38, 2015.

- 1265 Kozlova, E., Heimann, M., Zaehle, S., Worsey, J., Leppert, R., and Seifert, T.: Cape Verde Atmospheric Observatory: High-precision long-term atmospheric measurements of greenhouse gases (CO, CO₂, N₂O and CH₄) using Off-Axis Integrated-Cavity Output Spectroscopy (OA-ICOS) (2011 onwards), 2021.
- 1270 Labuschagne, C., Kuyper, B., Brunke, E.-G., Mokolo, T., van der Spuy, D., Martin, L., Mbambalala, E., Parker, B., Khan, M. A. H., Davies-Coleman, M. T., Shallcross, D. E., and Joubert, W.: A review of four decades of atmospheric trace gas measurements at Cape Point, South Africa, *Trans. R. Soc. South Afr.*, 73, 113–132, <https://doi.org/10.1080/0035919X.2018.1477854>, 2018.
- Lahoz, W. A. and Schneider, P.: Data assimilation: making sense of Earth Observation, *Front. Environ. Sci.*, 2, <https://doi.org/10.3389/fenvs.2014.00016>, 2014.
- 1275 Lambert, J.-C., Keppens, A., Compernelle, S., Eichmann, K.-U., de Graaf, M., Hubert, D., Langerock, B., Ludewig, A., Sha, M. K., Verhoelst, T., Wagner, T., Ahn, C., Argyrouli, A., Balis, D., Chan, K. L., Coldewey-Egbers, M., De Smedt, I., Eskes, H., Fjaeraa, A. M., Garane, K., Gleason, J. F., Goutail, F., Granville, J., Hedelt, P., Heue, K.-P., Jaross, G., Kleipool, Q., Koukouli, M. L., Lutz, R., Martinez Velarte, M. C., Michailidis, K., Pseftogkas, A., Nanda, S., Niemeijer, S., Pazmino, A., Pinardi, G., Richter, A., Rozemeijer, N., Sneep, M., Stein Zweers, D., Theys, N., Tilstra, G., Torres, O., Valks, P., van Geffen, J., Vigouroux, C., Wang, P., and Weber, M.: Quarterly Validation Report of the Copernicus Sentinel-5 Precursor Operational Data Products #24: April 2018 – August 2024, ESA/Copernicus S5P MPC, 2024.
- 1280 Landgraf, J.: Algorithm Theoretical Baseline Document: Carbon Monoxide Retrieval, 2019.
- Landgraf, J., aan de Brugh, J., Scheepmaker, R., Borsdorff, T., Hu, H., Houweling, S., Butz, A., Aben, I., and Hasekamp, O.: Carbon monoxide total column retrievals from TROPOMI shortwave infrared measurements, *Atmospheric Meas. Tech.*, 9, 4955–4975, <https://doi.org/10.5194/amt-9-4955-2016>, 2016.
- 1285 Laughner, J. L., Toon, G. C., Mendonca, J., Petri, C., Roche, S., Wunch, D., Blavier, J.-F., Griffith, D. W. T., Heikkinen, P., Keeling, R. F., Kiel, M., Kivi, R., Roehl, C. M., Stephens, B. B., Baier, B. C., Chen, H., Choi, Y., Deutscher, N. M., DiGangi, J. P., Gross, J., Herkommer, B., Jeseck, P., Laemmel, T., Lan, X., McGee, E., McKain, K., Miller, J., Morino, I., Notholt, J., Ohyama, H., Pollard, D. F., Rettinger, M., Riris, H., Rousogonous, C., Sha, M. K., Shiomi, K., Strong, K., Sussmann, R., Té, Y., Velasco, V. A., Wofsy, S. C., Zhou, M., and Wennberg, P. O.: The Total Carbon Column Observing Network’s GGG2020 data version, *Earth Syst. Sci. Data*, 16, 2197–2260, <https://doi.org/10.5194/essd-16-2197-2024>, 2024.
- 1290 Li, Y., Tong, D. Q., Ngan, F., Cohen, M. D., Stein, A. F., Kondragunta, S., Zhang, X., Ichoku, C., Hyer, E. J., and Kahn, R. A.: Ensemble PM_{2.5} Forecasting During the 2018 Camp Fire Event Using the HYSPLIT Transport and Dispersion Model, *J. Geophys. Res. Atmospheres*, 125, e2020JD032768, <https://doi.org/10.1029/2020JD032768>, 2020.
- 1295 Liu, Y., Kalnay, E., Zeng, N., Asrar, G., Chen, Z., and Jia, B.: Estimating surface carbon fluxes based on a local ensemble transform Kalman filter with a short assimilation window and a long observation window: an observing system simulation experiment test in GEOS-Chem 10.1, *Geosci. Model Dev.*, 12, 2899–2914, <https://doi.org/10.5194/gmd-12-2899-2019>, 2019.
- Liu, Y., Chen, J., Shi, Y., Zheng, W., Shan, T., and Wang, G.: Global Emissions Inventory from Open Biomass Burning (GEIOBB): utilizing Fengyun-3D global fire spot monitoring data, *Earth Syst. Sci. Data*, 16, 3495–3515, <https://doi.org/10.5194/essd-16-3495-2024>, 2024.
- 1300 Livings, D. M., Dance, S. L., and Nichols, N. K.: Unbiased ensemble square root filters, *Phys. Nonlinear Phenom.*, 237, 1021–1028, <https://doi.org/10.1016/j.physd.2008.01.005>, 2008.

- Lopez, J. P., Luo, M., Christensen, L. E., Loewenstein, M., Jost, H., Webster, C. R., and Osterman, G.: TES carbon monoxide validation during two AVE campaigns using the Argus and ALIAS instruments on NASA's WB-57F, *J. Geophys. Res. Atmospheres*, 113, <https://doi.org/10.1029/2007JD008811>, 2008.
- 1305 Lorente, A., Borsdorff, T., Butz, A., Hasekamp, O., aan de Brugh, J., Schneider, A., Wu, L., Hase, F., Kivi, R., Wunch, D., Pollard, D. F., Shiomi, K., Deutscher, N. M., Velasco, V. A., Roehl, C. M., Wennberg, P. O., Warneke, T., and Landgraf, J.: Methane retrieved from TROPOMI: improvement of the data product and validation of the first 2 years of measurements, *Atmospheric Meas. Tech.*, 14, 665–684, <https://doi.org/10.5194/amt-14-665-2021>, 2021.
- 1310 Lu, X., Jacob, D. J., Wang, H., Maasakkers, J. D., Zhang, Y., Scarpelli, T. R., Shen, L., Qu, Z., Sulprizio, M. P., Nesser, H., Bloom, A. A., Ma, S., Worden, J. R., Fan, S., Parker, R. J., Boesch, H., Gautam, R., Gordon, D., Moran, M. D., Reuland, F., Villasana, C. A. O., and Andrews, A.: Methane emissions in the United States, Canada, and Mexico: evaluation of national methane emission inventories and 2010–2017 sectoral trends by inverse analysis of in situ (GLOBALVIEWplus CH₄ ObsPack) and satellite (GOSAT) atmospheric observations, *Atmospheric Chem. Phys.*, 22, 395–418, <https://doi.org/10.5194/acp-22-395-2022>, 2022.
- 1315 Lutsch, E., Strong, K., Jones, D. B. A., Blumenstock, T., Conway, S., Fisher, J. A., Hannigan, J. W., Hase, F., Kasai, Y., Mahieu, E., Makarova, M., Morino, I., Nagahama, T., Notholt, J., Ortega, I., Palm, M., Poberovskii, A. V., Sussmann, R., and Warneke, T.: Detection and attribution of wildfire pollution in the Arctic and northern midlatitudes using a network of Fourier-transform infrared spectrometers and GEOS-Chem, *Atmospheric Chem. Phys.*, 20, 12813–12851, <https://doi.org/10.5194/acp-20-12813-2020>, 2020.
- 1320 Maasakkers, J. D., Jacob, D. J., Sulprizio, M. P., Scarpelli, T. R., Nesser, H., Sheng, J.-X., Zhang, Y., Hersher, M., Bloom, A. A., Bowman, K. W., Worden, J. R., Janssens-Maenhout, G., and Parker, R. J.: Global distribution of methane emissions, emission trends, and OH concentrations and trends inferred from an inversion of GOSAT satellite data for 2010–2015, *Atmospheric Chem. Phys.*, 19, 7859–7881, <https://doi.org/10.5194/acp-19-7859-2019>, 2019.
- 1325 Makarova, M., Poberovskii, A., Polyakov, A., Imkhasin, K. H., Ionov, D., Makarov, B., Kostsov, V., Foka, S., and Abakumov, E.: Trends of Key Greenhouse Gases as Measured in 2009–2022 at the FTIR Station of St. Petersburg State University, *Remote Sens.*, 16, 1996, <https://doi.org/10.3390/rs16111996>, 2024.
- Mataveli, G., Jones, M. W., Carmenta, R., Sanchez, A., Dutra, D. J., Chaves, M., de Oliveira, G., Anderson, L. O., and Aragão, L. E. O. C.: Deforestation falls but rise of wildfires continues degrading Brazilian Amazon forests, *Glob. Change Biol.*, 30, e17202, <https://doi.org/10.1111/gcb.17202>, 2024.
- 1330 McKain, K., Sweeney, C., Baier, B., Crotwell, A., Crotwell, M., Handley, P., Higgs, J., Legard, T., Madronich, M., Miller, J. B., Moglia, E., Mund, J., Newberger, T., Wolter, S., and NOAA Global Monitoring Laboratory.: NOAA Global Greenhouse Gas Reference Network Flask-Air PFP Sample Measurements of CO₂, CH₄, CO, N₂O, H₂, SF₆ and isotopic ratios collected from aircraft vertical profiles (2024-08-12), <https://doi.org/10.15138/39HR-9N34>, 2024.
- 1335 Ménard, R. and Chang, L.-P.: Assimilation of Stratospheric Chemical Tracer Observations Using a Kalman Filter. Part II: χ^2 -Validated Results and Analysis of Variance and Correlation Dynamics, *Mon. Weather Rev.*, 128, 2672–2686, [https://doi.org/10.1175/1520-0493\(2000\)128%253C2672:AOSCTO%253E2.0.CO;2](https://doi.org/10.1175/1520-0493(2000)128%253C2672:AOSCTO%253E2.0.CO;2), 2000.
- Migliorini, S.: Information-based data selection for ensemble data assimilation, *Q. J. R. Meteorol. Soc.*, 139, 2033–2054, <https://doi.org/10.1002/qj.2104>, 2013.
- 1340 Miyazaki, K., Eskes, H. J., Sudo, K., Takigawa, M., van Weele, M., and Boersma, K. F.: Simultaneous assimilation of satellite NO₂, O₃, CO, and HNO₃ data for the analysis of tropospheric chemical composition and emissions, *Atmospheric Chem. Phys.*, 12, 9545–9579, <https://doi.org/10.5194/acp-12-9545-2012>, 2012.

- Miyazaki, K., Eskes, H. J., and Sudo, K.: A tropospheric chemistry reanalysis for the years 2005–2012 based on an assimilation of OMI, MLS, TES, and MOPITT satellite data, *Atmospheric Chem. Phys.*, 15, 8315–8348, <https://doi.org/10.5194/acp-15-8315-2015>, 2015.
- 1345 Miyazaki, K., Bowman, K., Sekiya, T., Eskes, H., Boersma, F., Worden, H., Livesey, N., Payne, V. H., Sudo, K., Kanaya, Y., Takigawa, M., and Ogochi, K.: Updated tropospheric chemistry reanalysis and emission estimates, TCR-2, for 2005–2018, *Earth Syst. Sci. Data*, 12, 2223–2259, <https://doi.org/10.5194/essd-12-2223-2020>, 2020.
- Morino, I., Ohyama, H., Hori, A., and Ikegami, H.: TCCON data from Rikubetsu (JP), Release GGG2020.R0 (R0), <https://doi.org/10.14291/tcon.ggg2020.rikubetsu01.R0>, 2022a.
- 1350 Morino, I., Ohyama, H., Hori, A., and Ikegami, H.: TCCON data from Tsukuba (JP), 125HR, Release GGG2020.R0 (R0), <https://doi.org/10.14291/tcon.ggg2020.tsukuba02.R0>, 2022b.
- Nguyen, H. M., He, J., and Wooster, M. J.: Biomass burning CO, PM and fuel consumption per unit burned area estimates derived across Africa using geostationary SEVIRI fire radiative power and Sentinel-5P CO data, *Atmospheric Chem. Phys.*, 23, 2089–2118, <https://doi.org/10.5194/acp-23-2089-2023>, 2023.
- 1355 O'Neill, S. M., Diao, M., Raffuse, S., Al-Hamdan, M., Barik, M., Jia, Y., Reid, S., Zou, Y., Tong, D., West, J. J., Wilkins, J., Marsha, A., Freedman, F., Vargo, J., Larkin, N. K., Alvarado, E., and Loesche, P.: A multi-analysis approach for estimating regional health impacts from the 2017 Northern California wildfires, *J. Air Waste Manag. Assoc.*, 71, 791–814, <https://doi.org/10.1080/10962247.2021.1891994>, 2021.
- Palmer, P. I., Jacob, D. J., Jones, D. B. A., Heald, C. L., Yantosca, R. M., Logan, J. A., Sachse, G. W., and Streets, D. G.: Inverting for emissions of carbon monoxide from Asia using aircraft observations over the western Pacific, *J. Geophys. Res. Atmospheres*, 108, <https://doi.org/10.1029/2003JD003397>, 2003.
- 1360 Parker, H. A., Laughner, J. L., Toon, G. C., Wunch, D., Roehl, C. M., Iraci, L. T., Podolske, J. R., McKain, K., Baier, B. C., and Wennberg, P. O.: Inferring the vertical distribution of CO and CO₂ from TCCON total column values using the TARDISS algorithm, *ATMOSPHERIC Meas. Tech.*, 16, 2601–2625, <https://doi.org/10.5194/amt-16-2601-2023>, 2023.
- 1365 Patel, A., Mallik, C., Chandra, N., Patra, P. K., and Steinbacher, M.: Revisiting regional and seasonal variations in decadal carbon monoxide variability: Global reversal of growth rate, *Sci. Total Environ.*, 909, 168476, <https://doi.org/10.1016/j.scitotenv.2023.168476>, 2024.
- Pendergrass, D., Yantosca, B., and Voshtani, S.: drewpendergrass/CHEEREIO: CHEEREIO v1.3.1 release, , <https://doi.org/10.5281/zenodo.11534085>, 2024.
- 1370 Pendergrass, D. C., Jacob, D. J., Nesser, H., Varon, D. J., Sulprizio, M., Miyazaki, K., and Bowman, K. W.: CHEEREIO 1.0: a versatile and user-friendly ensemble-based chemical data assimilation and emissions inversion platform for the GEOS-Chem chemical transport model, *Geosci. Model Dev.*, 16, 4793–4810, <https://doi.org/10.5194/gmd-16-4793-2023>, 2023.
- Plant, G., Kort, E. A., Murray, L. T., Maasakkers, J. D., and Aben, I.: Evaluating urban methane emissions from space using TROPOMI methane and carbon monoxide observations, *Remote Sens. Environ.*, 268, 112756, <https://doi.org/10.1016/j.rse.2021.112756>, 2022.
- 1375 Pollard, D. F., Robinson, J., and Shiona, H.: TCCON data from Lauder (NZ), Release GGG2020.R0 (R0), <https://doi.org/10.14291/tcon.ggg2020.lauder03.R0>, 2022.

- Pommier, M., McLinden, C. A., and Deeter, M.: Relative changes in CO emissions over megacities based on observations from space, *Geophys. Res. Lett.*, 40, 3766–3771, <https://doi.org/10.1002/grl.50704>, 2013.
- 1380 Pope, R. J., Kerridge, B. J., Siddans, R., Latter, B. G., Chipperfield, M. P., Arnold, S. R., Ventress, L. J., Pimlott, M. A., Graham, A. M., Knappett, D. S., and Rigby, R.: Large Enhancements in Southern Hemisphere Satellite-Observed Trace Gases Due to the 2019/2020 Australian Wildfires, *J. Geophys. Res. Atmospheres*, 126, e2021JD034892, <https://doi.org/10.1029/2021JD034892>, 2021.
- 1385 Reza Koohkan, M. and Bocquet, M.: Accounting for representativeness errors in the inversion of atmospheric constituent emissions: application to the retrieval of regional carbon monoxide fluxes, *Tellus B Chem. Phys. Meteorol.*, 64, 19047, <https://doi.org/10.3402/tellusb.v64i0.19047>, 2012.
- Rodgers, C. D.: *Inverse Methods For Atmospheric Sounding: Theory And Practice*, World Scientific, 256 pp., 2000.
- Saikawa, E., Kim, H., Zhong, M., Avramov, A., Zhao, Y., Janssens-Maenhout, G., Kurokawa, J., Klimont, Z., Wagner, F., Naik, V., Horowitz, L. W., and Zhang, Q.: Comparison of emissions inventories of anthropogenic air pollutants and greenhouse gases in China, *Atmospheric Chem. Phys.*, 17, 6393–6421, <https://doi.org/10.5194/acp-17-6393-2017>, 2017.
- 1390 Schneising, O., Buchwitz, M., Reuter, M., Bovensmann, H., and Burrows, J. P.: Severe Californian wildfires in November 2018 observed from space: the carbon monoxide perspective, *Atmospheric Chem. Phys.*, 20, 3317–3332, <https://doi.org/10.5194/acp-20-3317-2020>, 2020.
- 1395 Schuldt, K. N., Aalto, T., Arlyn Andrews, Apadula, F., Arnold, S., Baier, B., Bani, L., Bergamaschi, P., Biermann, T., Biraud, S. C., Blot, R., Brand, W. A., Francescopiero Calzolari, Chen, G., Huilin Chen, Colomb, A., Commane, R., Condori, L., Conen, F., Conil, S., Couret, C., Cristofanelli, P., Cuevas, E., Daube, B., Davis, K. J., Delmotte, M., Dickerson, R., DiGangi, J. P., Diskin, G., Elsasser, M., Emmenegger, L., Fischer, M. L., Forster, G., Fuente-Lastra, M., Gehrlein, T., Gerbig, C., Hatakka, J., Heimann, M., Heliasz, M., Heltai, D., Hermans, C., Hermansen, O., Hoheisel, A., Holst, J., Di Iorio, T., Jaffe, D. A., Jordan, A., Karion, A., Kazan, V., Keronen, P., Kneuer, T., Kolari, P., Kominkova, K., Kort, E., Kozlova, E., Krummel, P. B., Kubistin, D., Kumps, N., Langenfelds, R. L., Lanza, A., Laurent, O., Laurila, T., Lavric, J., Lee, J., Lehner, I., Lehtinen, K., Leppert, R., 1400 Leskinen, A., Leuenberger, M., Levula, J., Lindauer, M., Lindroth, A., Mikael Ottosson Löfvenius, Loh, Z. M., Lopez, M., Lunder, C. R., Mammarella, I., Manca, G., Marek, M. V., Marklund, P., Martin, M. Y., De Mazière, M., McKain, K., Meinhardt, F., Jean-Marc Metzger, Miles, N. L., Miller, C. E., Miller, J. B., Mölder, M., Moossen, H., Müller-Williams, J., Myhre, C. L., Nédélec, P., Obersteiner, F., Peltola, O., Petron, G., Piacentino, S., Pichon, J. M., Pickers, P., et al.: Multi-laboratory compilation of atmospheric carbon monoxide data for the period 1989-2022; 1405 obspack_co_1_GLOBALVIEWplus_v4.0_2024-02-13, <https://doi.org/10.25925/20231220>, 2024.
- Sha, M. K., Langerock, B., Blavier, J.-F. L., Blumenstock, T., Borsdorff, T., Buschmann, M., Dehn, A., De Mazière, M., Deutscher, N. M., Feist, D. G., García, O. E., Griffith, D. W. T., Grutter, M., Hannigan, J. W., Hase, F., Heikkinen, P., Hermans, C., Iraci, L. T., Jeseck, P., Jones, N., Kivi, R., Kumps, N., Landgraf, J., Lorente, A., Mahieu, E., Makarova, M. V., Mellqvist, J., Metzger, J.-M., Morino, I., Nagahama, T., Notholt, J., Ohyama, H., Ortega, I., Palm, M., Petri, C., Pollard, D. F., Rettinger, 1410 M., Robinson, J., Roche, S., Roehl, C. M., Röhl, A. N., Rousogonous, C., Schneider, M., Shiomi, K., Smale, D., Stremme, W., Strong, K., Sussmann, R., Té, Y., Uchino, O., Velasco, V. A., Vigouroux, C., Vrekoussis, M., Wang, P., Warneke, T., Wizenberg, T., Wunch, D., Yamanouchi, S., Yang, Y., and Zhou, M.: Validation of methane and carbon monoxide from Sentinel-5 Precursor using TCCON and NDACC-IRWG stations, *Atmospheric Meas. Tech.*, 14, 6249–6304, <https://doi.org/10.5194/amt-14-6249-2021>, 2021.
- 1415 Shahrokhi, N., Rayner, P. J., Silver, J. D., and Thomas, S.: Urban-scale variational flux inversion for CO Using TROPOMI total-column retrievals: A case study of Tehran, *Atmos. Environ.*, 311, 120009, <https://doi.org/10.1016/j.atmosenv.2023.120009>, 2023.

- Smale, D., Strahan, S. E., Querel, R., Frieß, U., Nedoluha, G. E., Nichol, S. E., Robinson, J., Boyd, I., Kotkamp, M., Gomez, R. M., Murphy, M., Tran, H., and McGaw, J.: Evolution of observed ozone, trace gases, and meteorological variables over Arrival Heights, Antarctica (77.8°S, 166.7°E) during the 2019 Antarctic stratospheric sudden warming, *Tellus B Chem. Phys. Meteorol.*, 73, 1–18, <https://doi.org/10.1080/16000889.2021.1933783>, 2021.
- Sokolik, I. N., Soja, A. J., DeMott, P. J., and Winker, D.: Progress and Challenges in Quantifying Wildfire Smoke Emissions, Their Properties, Transport, and Atmospheric Impacts, *J. Geophys. Res. Atmospheres*, 124, 13005–13025, <https://doi.org/10.1029/2018JD029878>, 2019.
- Stephens, B. B., Gurney, K. R., Tans, P. P., Sweeney, C., Peters, W., Bruhwiler, L., Ciais, P., Ramonet, M., Bousquet, P., Nakazawa, T., Aoki, S., Machida, T., Inoue, G., Vinnichenko, N., Lloyd, J., Jordan, A., Heimann, M., Shibistova, O., Langenfelds, R. L., Steele, L. P., Francey, R. J., and Denning, A. S.: Weak Northern and Strong Tropical Land Carbon Uptake from Vertical Profiles of Atmospheric CO₂, *Science*, 316, 1732–1735, <https://doi.org/10.1126/science.1137004>, 2007.
- Stockwell, C. E., Bela, M. M., Coggon, M. M., Gkatzelis, G. I., Wiggins, E., Gargulinski, E. M., Shingler, T., Fenn, M., Griffin, D., Holmes, C. D., Ye, X., Saide, P. E., Bourgeois, I., Peischl, J., Womack, C. C., Washenfelder, R. A., Veres, P. R., Neuman, J. A., Gilman, J. B., Lamplugh, A., Schwantes, R. H., McKeen, S. A., Wisthaler, A., Piel, F., Guo, H., Campuzano-Jost, P., Jimenez, J. L., Fried, A., Hanisco, T. F., Huey, L. G., Perring, A., Katich, J. M., Diskin, G. S., Nowak, J. B., Bui, T. P., Halliday, H. S., DiGangi, J. P., Pereira, G., James, E. P., Ahmadov, R., McLinden, C. A., Soja, A. J., Moore, R. H., Hair, J. W., and Warneke, C.: Airborne Emission Rate Measurements Validate Remote Sensing Observations and Emission Inventories of Western U.S. Wildfires, *Environ. Sci. Technol.*, 56, 7564–7577, <https://doi.org/10.1021/acs.est.1c07121>, 2022.
- Sussmann, R. and Rettinger, M.: TCCON data from Garmisch (DE), Release GGG2020.R0 (R0), <https://doi.org/10.14291/tcon.ggg2020.garmisch01.R0>, 2023.
- Sweeney, C., Karion, A., Wolter, S., Newberger, T., Guenther, D., Higgs, J. A., Andrews, A. E., Lang, P. M., Neff, D., Dlugokencky, E., Miller, J. B., Montzka, S. A., Miller, B. R., Masarie, K. A., Biraud, S. C., Novelli, P. C., Crotwell, M., Crotwell, A. M., Thoning, K., and Tans, P. P.: Seasonal climatology of CO₂ across North America from aircraft measurements in the NOAA/ESRL Global Greenhouse Gas Reference Network, *J. Geophys. Res. Atmospheres*, 120, 5155–5190, <https://doi.org/10.1002/2014JD022591>, 2015.
- Takatsuji, S.: Atmospheric CO at Minamitorishima by Japan Meteorological Agency (2024-05-16-0623), https://doi.org/10.50849/WDCGG_0001-2029-3001-01-01-9999, 2024a.
- Takatsuji, S.: Atmospheric CO at Ryori by Japan Meteorological Agency (2024-05-16-0623), https://doi.org/10.50849/WDCGG_0001-2012-3001-01-01-9999, 2024b.
- Takatsuji, S.: Atmospheric CO at Yonagunijima by Japan Meteorological Agency (2024-05-16-0623), https://doi.org/10.50849/WDCGG_0001-2028-3001-01-01-9999, 2024c.
- Tang, W., Emmons, L. K., Jr, A. F. A., Gaubert, B., Knote, C., Tilmes, S., Buchholz, R. R., Pfister, G. G., Diskin, G. S., Blake, D. R., Blake, N. J., Meinardi, S., DiGangi, J. P., Choi, Y., Woo, J.-H., He, C., Schroeder, J. R., Suh, I., Lee, H.-J., Jo, H.-Y., Kanaya, Y., Jung, J., Lee, Y., and Kim, D.: Source Contributions to Carbon Monoxide Concentrations During KORUS-AQ Based on CAM-chem Model Applications, *J. Geophys. Res. Atmospheres*, 124, 2796–2822, <https://doi.org/10.1029/2018JD029151>, 2019.
- Tang, W., Gaubert, B., Emmons, L., Ziskin, D., Mao, D., Edwards, D., Arellano, A., Raeder, K., Anderson, J., and Worden, H.: Advantages of assimilating multispectral satellite retrievals of atmospheric composition: a demonstration using MOPITT carbon monoxide products, *Atmospheric Meas. Tech.*, 17, 1941–1963, <https://doi.org/10.5194/amt-17-1941-2024>, 2024.

- Tang, X., Zhu, J., Wang, Z. F., Wang, M., Gbaguidi, A., Li, J., Shao, M., Tang, G. Q., and Ji, D. S.: Inversion of CO emissions over Beijing and its surrounding areas with ensemble Kalman filter, *Atmos. Environ.*, 81, 676–686, <https://doi.org/10.1016/j.atmosenv.2013.08.051>, 2013.
- 1460 Tang, Z., Chen, J., and Jiang, Z.: Discrepancy in assimilated atmospheric CO over East Asia in 2015–2020 by assimilating satellite and surface CO measurements, *Atmospheric Chem. Phys.*, 22, 7815–7826, <https://doi.org/10.5194/acp-22-7815-2022>, 2022.
- 1465 Tang, Z., Jiang, Z., Chen, J., Yang, P., and Shen, Y.: The capabilities of the adjoint of GEOS-Chem model to support HEMCO emission inventories and MERRA-2 meteorological data, *Geosci. Model Dev.*, 16, 6377–6392, <https://doi.org/10.5194/gmd-16-6377-2023>, 2023.
- Tanimoto, H., Sawa, Y., Yonemura, S., Yumimoto, K., Matsueda, H., Uno, I., Hayasaka, T., Mukai, H., Tohjima, Y., Tsuboi, K., and Zhang, L.: Diagnosing recent CO emissions and ozone evolution in East Asia using coordinated surface observations, adjoint inverse modeling, and MOPITT satellite data, *Atmospheric Chem. Phys.*, 8, 3867–3880, <https://doi.org/10.5194/acp-8-3867-2008>, 2008.
- 1470 Té, Y., Jeseck, P., and Janssen, C.: TCCON data from Paris (FR), Release GGG2020.R0 (R0), <https://doi.org/10.14291/tcon.ggg2020.paris01.R0>, 2022.
- The International GEOS-Chem User Community: geoschem/GCClassic: GEOS-Chem Classic 14.1.1, , <https://doi.org/10.5281/zenodo.7696651>, 2023.
- 1475 Total Carbon Column Observing Network (TCCON) Team.: 2020 TCCON Data Release (GGG2020), <https://doi.org/10.14291/TCCON.GGG2020>, 2022.
- Travis, K. R., Jacob, D. J., Fisher, J. A., Kim, P. S., Marais, E. A., Zhu, L., Yu, K., Miller, C. C., Yantosca, R. M., Sulprizio, M. P., Thompson, A. M., Wennberg, P. O., Crounse, J. D., St. Clair, J. M., Cohen, R. C., Laughner, J. L., Dibb, J. E., Hall, S. R., Ullmann, K., Wolfe, G. M., Pollack, I. B., Peischl, J., Neuman, J. A., and Zhou, X.: Why do models overestimate surface ozone in the Southeast United States?, *Atmospheric Chem. Phys.*, 16, 13561–13577, <https://doi.org/10.5194/acp-16-13561-2016>, 2016.
- 1480 Turner, A. J., Jacob, D. J., Wecht, K. J., Maasakkers, J. D., Lundgren, E., Andrews, A. E., Biraud, S. C., Boesch, H., Bowman, K. W., Deutscher, N. M., Dubey, M. K., Griffith, D. W. T., Hase, F., Kuze, A., Notholt, J., Ohyama, H., Parker, R., Payne, V. H., Sussmann, R., Sweeney, C., Velasco, V. A., Warneke, T., Wennberg, P. O., and Wunch, D.: Estimating global and North American methane emissions with high spatial resolution using GOSAT satellite data, *Atmospheric Chem. Phys.*, 15, 7049–7069, <https://doi.org/10.5194/acp-15-7049-2015>, 2015.
- 1485 Turquety, S., Hadji-Lazaro, J., Clerbaux, C., Hauglustaine, D. A., Clough, S. A., Cassé, V., Schlüssel, P., and Mégie, G.: Operational trace gas retrieval algorithm for the Infrared Atmospheric Sounding Interferometer, *J. Geophys. Res. Atmospheres*, 109, <https://doi.org/10.1029/2004JD004821>, 2004.
- 1490 Varon, D. J., Jacob, D. J., Sulprizio, M., Estrada, L. A., Downs, W. B., Shen, L., Hancock, S. E., Nesser, H., Qu, Z., Penn, E., Chen, Z., Lu, X., Lorente, A., Tewari, A., and Randles, C. A.: Integrated Methane Inversion (IMI 1.0): a user-friendly, cloud-based facility for inferring high-resolution methane emissions from TROPOMI satellite observations, *Geosci. Model Dev.*, 15, 5787–5805, <https://doi.org/10.5194/gmd-15-5787-2022>, 2022.
- 1495 Veefkind, J. P., Aben, I., McMullan, K., Förster, H., de Vries, J., Otter, G., Claas, J., Eskes, H. J., de Haan, J. F., Kleipool, Q., van Weele, M., Hasekamp, O., Hoogeveen, R., Landgraf, J., Snel, R., Tol, P., Ingmann, P., Voors, R., Kruizinga, B., Vink, R., Visser, H., and Levelt, P. F.: TROPOMI on the ESA Sentinel-5 Precursor: A GMES mission for global observations of the

- atmospheric composition for climate, air quality and ozone layer applications, *Remote Sens. Environ.*, 120, 70–83, <https://doi.org/10.1016/j.rse.2011.09.027>, 2012.
- Voshtani, S., Ménard, R., Walker, T. W., and Hakami, A.: Assimilation of GOSAT Methane in the Hemispheric CMAQ; Part II: Results Using Optimal Error Statistics, *Remote Sens.*, 14, 375, <https://doi.org/10.3390/rs14020375>, 2022.
- 1500 Voshtani, S., Ménard, R., Walker, T. W., and Hakami, A.: Use of Assimilation Analysis in 4D-Var Source Inversion: Observing System Simulation Experiments (OSSEs) with GOSAT Methane and Hemispheric CMAQ, *Atmosphere*, 14, 758, <https://doi.org/10.3390/atmos14040758>, 2023.
- Wan, N., Xiong, X., Kluitenberg, G. J., Hutchinson, J. M. S., Aiken, R., Zhao, H., and Lin, X.: Estimation of biomass burning emission of NO₂ and CO from 2019–2020 Australia fires based on satellite observations, *Atmospheric Chem. Phys.*, 23, 711–724, <https://doi.org/10.5194/acp-23-711-2023>, 2023.
- 1505 Wang, J. S., Kawa, S. R., Collatz, G. J., Sasakawa, M., Gatti, L. V., Machida, T., Liu, Y., and Manyin, M. E.: A global synthesis inversion analysis of recent variability in CO₂ fluxes using GOSAT and in situ observations, *Atmospheric Chem. Phys.*, 18, 11097–11124, <https://doi.org/10.5194/acp-18-11097-2018>, 2018.
- Warner, J., Comer, M. M., Barnet, C. D., McMillan, W. W., Wolf, W., Maddy, E., and Sachse, G.: A comparison of satellite tropospheric carbon monoxide measurements from AIRS and MOPITT during INTEX-A, *J. Geophys. Res. Atmospheres*, 112, <https://doi.org/10.1029/2006JD007925>, 2007.
- 1510 Warner, J. X., Wei, Z., Strow, L. L., Barnet, C. D., Sparling, L. C., Diskin, G., and Sachse, G.: Improved agreement of AIRS tropospheric carbon monoxide products with other EOS sensors using optimal estimation retrievals, *Atmospheric Chem. Phys.*, 10, 9521–9533, <https://doi.org/10.5194/acp-10-9521-2010>, 2010.
- 1515 Weidmann, D., Brownsword, R., and Doniki, S.: TCCON data from Harwell, Oxfordshire (UK), Release GGG2020.R0 (R0), <https://doi.org/10.14291/tccon.ggg2020.harwell01.R0>, 2023.
- Wennberg, P. O., Roehl, C. M., Wunch, D., Blavier, J.-F., Toon, G. C., Allen, N. T., Treffers, R., and Laughner, J.: TCCON data from Caltech (US), Release GGG2020.R0 (R0), <https://doi.org/10.14291/tccon.ggg2020.pasadena01.R0>, 2022a.
- Wennberg, P. O., Wunch, D., Roehl, C. M., Blavier, J.-F., Toon, G. C., and Allen, N. T.: TCCON data from Lamont (US), Release GGG2020.R0 (R0), <https://doi.org/10.14291/tccon.ggg2020.lamont01.R0>, 2022b.
- 1520 Wennberg, P. O., Roehl, C. M., Wunch, D., Toon, G. C., Blavier, J.-F., Washenfelder, R., Keppel-Aleks, G., and Allen, N. T.: TCCON data from Park Falls (US), Release GGG2020.R1 (R1), <https://doi.org/10.14291/tccon.ggg2020.parkfalls01.R1>, 2022c.
- van der Werf, G. R., Randerson, J. T., Giglio, L., van Leeuwen, T. T., Chen, Y., Rogers, B. M., Mu, M., van Marle, M. J. E., Morton, D. C., Collatz, G. J., Yokelson, R. J., and Kasibhatla, P. S.: Global fire emissions estimates during 1997–2016, *Earth Syst. Sci. Data*, 9, 697–720, <https://doi.org/10.5194/essd-9-697-2017>, 2017.
- 1525 Wiedinmyer, C., Kimura, Y., McDonald-Buller, E. C., Emmons, L. K., Buchholz, R. R., Tang, W., Seto, K., Joseph, M. B., Barsanti, K. C., Carlton, A. G., and Yokelson, R.: The Fire Inventory from NCAR version 2.5: an updated global fire emissions model for climate and chemistry applications, *Geosci. Model Dev.*, 16, 3873–3891, <https://doi.org/10.5194/gmd-16-3873-2023>, 2023.
- 1530 WRAP: 2002 Fire Emission Inventory for the WRAP Region – Phase II, 2005.

- Wu, D., Liu, J., Wennberg, P. O., Palmer, P. I., Nelson, R. R., Kiel, M., and Eldering, A.: Towards sector-based attribution using intra-city variations in satellite-based emission ratios between CO₂ and CO, *Atmospheric Chem. Phys.*, 22, 14547–14570, <https://doi.org/10.5194/acp-22-14547-2022>, 2022.
- 1535 Wunch, D., Toon, G. C., Wennberg, P. O., Wofsy, S. C., Stephens, B. B., Fischer, M. L., Uchino, O., Abshire, J. B., Bernath, P., Biraud, S. C., Blavier, J.-F. L., Boone, C., Bowman, K. P., Browell, E. V., Campos, T., Connor, B. J., Daube, B. C., Deutscher, N. M., Diao, M., Elkins, J. W., Gerbig, C., Gottlieb, E., Griffith, D. W. T., Hurst, D. F., Jiménez, R., Keppel-Aleks, G., Kort, E. A., Macatangay, R., Machida, T., Matsueda, H., Moore, F., Morino, I., Park, S., Robinson, J., Roehl, C. M., Sawa, Y., Sherlock, V., Sweeney, C., Tanaka, T., and Zondlo, M. A.: Calibration of the Total Carbon Column Observing Network using aircraft profile data, *Atmospheric Meas. Tech.*, 3, 1351–1362, <https://doi.org/10.5194/amt-3-1351-2010>, 2010.
- 1540 Wunch, D., Toon, G. C., Blavier, J.-F. L., Washenfelder, R. A., Notholt, J., Connor, B. J., Griffith, D. W. T., Sherlock, V., and Wennberg, P. O.: The Total Carbon Column Observing Network, *Philos. Trans. R. Soc. Math. Phys. Eng. Sci.*, 369, 2087–2112, <https://doi.org/10.1098/rsta.2010.0240>, 2011.
- 1545 Wunch, D., Jones, D. B. A., Toon, G. C., Deutscher, N. M., Hase, F., Notholt, J., Sussmann, R., Warneke, T., Kuenen, J., Denier van der Gon, H., Fisher, J. A., and Maasakkers, J. D.: Emissions of methane in Europe inferred by total column measurements, *Atmospheric Chem. Phys.*, 19, 3963–3980, <https://doi.org/10.5194/acp-19-3963-2019>, 2019.
- Wunch, D., Mendonca, J., Colebatch, O., Allen, N. T., Blavier, J.-F., Kunz, K., Roche, S., Hedelius, J., Neufeld, G., Springett, S., Worthy, D., Kessler, R., and Strong, K.: TCCON data from East Trout Lake, SK (CA), Release GGG2020.R0 (R0), <https://doi.org/10.14291/tcon.ggg2020.easttroutlake01.R0>, 2022.
- 1550 Wunch, D., Laughner, J., Toon, G. C., Roehl, C. M., Wennberg, P. O., Mill'an, L. F., Deutscher, N. M., Warneke, T., Pollard, D. F., Feist, D. G., Strong, K., McGee, E., Roche, S., Mendonca, J., Kivi, R., Heikkinen, P., Hase, F., Sha, M. K., De Mazière, M., Sussmann, R., Rettinger, M., Pak, N. M., Morino, I., Velasco, V. A., Griffith, D. W. T., Notholt, J., Petri, C., Buschmann, M., Hachmeister, J., Doniki, S., Weidmann, D., Rousogonous, C., Vrekoussis, M., Ohyama, H., Oh, Y.-S., García, O. E., Robinson, J., Dubey, M., Mingqiang, Z., Wang, P., Té, Y., Jeseck, P., Iraci, L., Podolske, J., Shiomi, K., and Kawakami, S.: The Total Carbon Column Observing Network's GGG2020 Data Version: Data Quality, Comparison with GGG2014, and Future Outlook, 2025.
- 1555 Yang, Z., Washenfelder, R. A., Keppel-Aleks, G., Krakauer, N. Y., Randerson, J. T., Tans, P. P., Sweeney, C., and Wennberg, P. O.: New constraints on Northern Hemisphere growing season net flux, *Geophys. Res. Lett.*, 34, <https://doi.org/10.1029/2007GL029742>, 2007.
- 1560 You, C. and Xu, C.: Delayed wildfires in 2020 promote snowpack melting in the western United States, *Proc. Natl. Acad. Sci.*, 120, e2218087120, <https://doi.org/10.1073/pnas.2218087120>, 2023.
- Yumimoto, K. and Uno, I.: Adjoint inverse modeling of CO emissions over Eastern Asia using four-dimensional variational data assimilation, *Atmos. Environ.*, 40, 6836–6845, <https://doi.org/10.1016/j.atmosenv.2006.05.042>, 2006.
- 1565 Žagar, N., Anderson, J., Collins, N., Hoar, T., Raeder, K., Lei, L., and Tribbia, J.: Scale-Dependent Representation of the Information Content of Observations in the Global Ensemble Kalman Filter Data Assimilation, *Mon. Weather Rev.*, 144, 2927–2945, <https://doi.org/10.1175/MWR-D-15-0401.1>, 2016.
- Zander, R., Mahieu, E., Demoulin, P., Duchatelet, P., Roland, G., Servais, C., Mazière, M. D., Reimann, S., and Rinsland, C. P.: Our changing atmosphere: Evidence based on long-term infrared solar observations at the Jungfraujoch since 1950, *Sci. Total Environ.*, 391, 184–195, <https://doi.org/10.1016/j.scitotenv.2007.10.018>, 2008.

- 1570 Zhang, L., Montuoro, R., McKeen, S. A., Baker, B., Bhattacharjee, P. S., Grell, G. A., Henderson, J., Pan, L., Frost, G. J., McQueen, J., Saylor, R., Li, H., Ahmadov, R., Wang, J., Stajner, I., Kondragunta, S., Zhang, X., and Li, F.: Development and evaluation of the Aerosol Forecast Member in the National Center for Environment Prediction (NCEP)'s Global Ensemble Forecast System (GEFS-Aerosols v1), *Geosci. Model Dev.*, 15, 5337–5369, <https://doi.org/10.5194/gmd-15-5337-2022>, 2022.
- 1575 Zhang, X., Kondragunta, S., Ram, J., Schmidt, C., and Huang, H.-C.: Near-real-time global biomass burning emissions product from geostationary satellite constellation, *J. Geophys. Res. Atmospheres*, 117, <https://doi.org/10.1029/2012JD017459>, 2012.
- Zhang, X., Kondragunta, S., Silva, A. D., Lu, S., Ding, H., Li, F., and Zhu, Y.: TITLE: THE BLENDED GLOBAL BIOMASS BURNING EMISSIONS PRODUCT FROM MODIS AND VIIRS OBSERVATIONS (GBBEPx) VERSION 3., n.d.
- 1580 Zhang, Z., Zhang, L., Xu, H., Creed, I. F., Blanco, J. A., Wei, X., Sun, G., Asbjornsen, H., and Bishop, K.: Forest water-use efficiency: Effects of climate change and management on the coupling of carbon and water processes, *For. Ecol. Manag.*, 534, 120853, <https://doi.org/10.1016/j.foreco.2023.120853>, 2023.
- Zhao, J., Ciais, P., Chevallier, F., Canadell, J. G., van der Velde, I. R., Chuvieco, E., Chen, Y., Zhang, Q., He, K., and Zheng, B.: Enhanced CH₄ emissions from global wildfires likely due to undetected small fires, *Nat. Commun.*, 16, 804, <https://doi.org/10.1038/s41467-025-56218-w>, 2025.
- 1585 Zheng, B., Chevallier, F., Yin, Y., Ciais, P., Fortems-Cheiney, A., Deeter, M. N., Parker, R. J., Wang, Y., Worden, H. M., and Zhao, Y.: Global atmospheric carbon monoxide budget 2000–2017 inferred from multi-species atmospheric inversions, *Earth Syst. Sci. Data*, 11, 1411–1436, <https://doi.org/10.5194/essd-11-1411-2019>, 2019.
- Zupanski, D., Hou, A. Y., Zhang, S. Q., Zupanski, M., Kummerow, C. D., and Cheung, S. H.: Applications of information theory in ensemble data assimilation, *Q. J. R. Meteorol. Soc.*, 133, 1533–1545, <https://doi.org/10.1002/qj.123>, 2007.
- 1590 Zupanski, M.: Maximum Likelihood Ensemble Filter: Theoretical Aspects, *Mon. Weather Rev.*, 133, 1710–1726, <https://doi.org/10.1175/MWR2946.1>, 2005.