

Towards an improved understanding of the impact of clouds and precipitation on the representation of aerosols over the Boreal Forest in GCMs

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Abstract

Global climate models (GCMs) face uncertainties in estimating Earth's radiative budget due to aerosol-cloud interactions (ACI). Accurate particle number size distributions (PNSDs) are crucial for improving ACI representation, requiring precise modelling of aerosol sources and sinks. Using a Lagrangian trajectory framework, we examine how clouds and precipitation influence aerosols during transport, and thereby influence aerosol–cloud relationships in the boreal forest. Two GCMs, the United Kingdom Earth System Model (UKESM1) and ECHAM6.3-HAM2.3-MOZ1.0 with the SALSA2.0 aerosol module (ECHAM-SALSA), are complemented with model-derived trajectories and evaluated against in-situ observations, which are accompanied by reanalysis trajectories. Overall aerosol–precipitation trends are similar between GCMs and observations. However, seasonal differences emerge: in summer, UKESM1 exhibits more efficient aerosol removal via precipitation than ECHAM-SALSA and observations, whereas in winter, the opposite is observed. These differences coincide with key variables controlling aerosol activation, such as sub-grid scale updraught velocities and PNSDs. For example, in winter, removal of total aerosol mass in ECHAM-SALSA was stronger than in UKESM1, coinciding with higher activated fractions and larger sub-grid scale updraught velocities in ECHAM-SALSA. For both GCMs, cloud processing along trajectories increased SO₄ mass, mainly in the accumulation mode, consistent with observations and model parametrizations. Discrepancies arise more from differences in PNSDs and updraught velocities

46 than from wet removal parametrizations, an example being the underrepresentation of small particles in UKESM1. While
47 our findings are representative of boreal region with predominantly stratiform precipitation, further work is needed to
48 evaluate their applicability to other regions.

49 **1 Introduction**

50 Atmospheric aerosol particle concentrations are influenced by their sources and sinks which affect their lifetimes in the
51 atmosphere, and also play a significant role in our climate system through different mechanisms. One of the most
52 important mechanisms are aerosol-cloud interactions (ACI), which are still causing the largest uncertainties on the effects
53 of aerosols on Earth's radiative budget in global climate models (GCMs, Boucher, 2013; Watson-Parris et al., 2019;
54 Bellouin et al., 2020; Forster et al., 2021), and therefore partly masking the warming effect by greenhouse gases (Bauer
55 et al., 2022; Quaas et al., 2022). It is critical, therefore, that the microphysical processes influencing ACIs are well
56 understood and accurately modelled. To accurately simulate ACI in GCMs, the aerosol number size distributions need to
57 be correctly described (e.g., Mann et al., 2010). Traditionally, discrepancies in particle size distributions between
58 observations and models exceed those between modal and sectional approaches, with sectional methods dividing the
59 distribution into discrete size bins (Mann et al., 2012). However, larger differences in concentrations may emerge when
60 chemistry of the aerosols is inspected (Laakso et al., 2022). On the other hand, to accurately represent the aerosol number
61 size distributions, GCMs also need to accurately represent the source and sink processes that act on the aerosol during its
62 lifetime and transport in the atmosphere. The impact of precipitation on the evolution of the size distribution is very
63 important (e.g., Browse et al., 2014; Khadir et al., 2023), but remains a major uncertainty in the GCMs. Often, when
64 GCM parametrizations are assessed the models are evaluated against observations or other GCMs by inspecting
65 differences in averages of variables (or relationships between multiple variables) over certain time spans (e.g., Blichner
66 et al., 2024; Gliß et al., 2021; Labe and Barnes, 2022; Maher et al., 2021; Pathak et al., 2023) in a Eulerian perspective.
67 However, GCM evaluations in which the evolution of aerosols and other variables is followed over both time and space
68 in more detail using GCM Lagrangian trajectory-based evaluation frameworks that have been recently introduced (e.g.,
69 Kim et al., 2020). Such frameworks pave the way for the development of more rigorous observational constraints on
70 uncertain physical and chemical aerosol processes for GCM evaluation, by including temporal and spatial information
71 associated with the air-mass history.

72 ACIs include scavenging of aerosol particles by precipitation, cloud droplets and ice crystals. Wet scavenging is one of
73 the most efficient removal routes of particles from the atmosphere (e.g., Ohata et al., 2016; Liu et al., 2020). Wet
74 scavenging of aerosol particles can be further divided into in-cloud scavenging and below cloud scavenging. Wet
75 scavenging via in-cloud scavenging involves the loss of aerosol particles when they become activated into cloud droplets
76 or ice crystals (nucleation scavenging) which can then further collide with interstitial aerosols in-cloud (e.g., Ohata et al.,
77 2016; Seinfeld and Pandis, 2016). Below-cloud scavenging concerns the removal of aerosol by rainfall from the collection
78 of particles due to collisions with falling raindrops and snow and ice from precipitation (e.g., Ohata et al., 2016). Current
79 understanding identifies the contribution of in-cloud scavenging, followed by removal via precipitation to be, on average,
80 the most important sink globally for accumulation mode particles (particle diameter $d_p \sim 100\text{-}1000\text{ nm}$). Ultrafine ($d_p <$
81 100 nm) and coarse particles ($d_p > 1\text{ }\mu\text{m}$), on the other hand, are more efficiently removed by below-cloud scavenging
82 (e.g., Andronache, 2003; Textor et al., 2006; Croft et al., 2009; Ohata et al., 2016). In addition to wet scavenging, clouds
83 can also alter the particle properties through aqueous phase oxidation processes. For example, sulfate production due to
84 oxidation of gaseous sulfur dioxide inside clouds is considered as one of the most important mass addition processes for
85 sulfate (e.g., Ervens, 2015 and references therein). Production of organics through aqueous phase processes has also been
86 reported in some environments (e.g., Ervens et al., 2018; Lamkaddam et al., 2021).

Investigation of the effects of precipitation and clouds has traditionally been Eulerian, in which local estimates of precipitation are employed (e.g., Wang et al., 2021). Lagrangian approaches, in which air mass trajectories are exploited to examine the effects of precipitation on aerosols and their composition as the air masses travel to the receptor location, have, however, increased in popularity during the recent years (Dadashazar et al., 2021; Heslin-Rees et al., 2024; Isokääntä et al., 2022; Kesti et al., 2020; Khadir et al., 2023; Tunved et al., 2004, 2013; Tunved and Ström, 2019). These types of studies can provide significantly more detailed insights by considering the interplay between aerosols, clouds and precipitation during air mass history, that cannot be achieved using Eulerian approaches. All these studies investigated how the total accumulated precipitation experienced along air-mass trajectories derived from reanalysis data affects a particle size distribution measured at a specific receptor site. Tunved et al. (2013), for example, investigated aerosols in the Arctic (Zeppelin station, Ny-Ålesund, Norway) and observed strong removal of sub-micron particulate mass up to 10 mm of accumulated precipitation. They suggested the in-cloud scavenging (followed by removal via precipitation) is the dominant removal pathway, as larger particles showed first a decrease in their concentration as a function of accumulated precipitation during transport, followed by the removal smaller sizes. Kesti et al. (2020) studied aerosols at the humid tropical monsoon climate in the Maldives, and observed more efficient removal of the accumulation mode particles with increasing accumulated precipitation, when compared to the smaller particle sizes. Dadashazar et al. (2021) studied sub-tropical environments in Bermuda and concluded that PM_{2.5} mass experienced the strongest sensitivity to accumulated precipitation up to 5 mm whereas precipitation exceeding this limit had no major effects on the particulate mass. Khadir et al. (2023) further reported that precipitation can, in some instances, serve as a source of aerosols.

In addition to the effects of precipitation for aerosols, a previous study by Isokääntä et al. (2022) used relative humidity (>94%) as a proxy for in-cloud exposure in boreal air masses and found a pronounced increase in sulfate mass in air masses recently influenced by non-precipitating clouds, while no significant aqueous-phase production of organic aerosol was observed—likely due to dominant gas-phase biogenic sources. This is consistent with findings from central Sweden (Graham et al., 2020). These earlier results suggest that sulfate may be more strongly affected by cloud processing and wet removal than organic aerosol, with removal efficiency likely influenced by factors such as precipitation timing, aerosol type, and the stage of the air mass trajectory. Our study builds on this by exploring these aspects across multiple models and observations, employing the GCM Lagrangian evaluation framework presented by Kim et al. (2020). With this framework air mass trajectories can be obtained from global GCM simulations. This is achieved by co-locating multiple variables (for example, aerosol size distribution and chemical composition) from the GCMs to air mass trajectories calculated from the GCM meteorological data (Kim et al., 2020). This methodology allows us to transparently evaluate and compare the wet scavenging and aqueous-phase processing between the observations and GCMs within the Lagrangian trajectory framework in unprecedented detail.

This study compares the effects of wet processing (wet removal and aqueous-phase processing) on modelled aerosol size distributions with long-term observations from Hyytiälä, Finland. Observational trajectories are based on ERA-Interim reanalysis, while model trajectories are calculated using meteorology data from GCM AMIP-style simulations in which wind fields were nudged to ERA-Interim. The GCMs used in this study include UKESM1 (United Kingdom Earth System Model, e.g., Sellar et al., 2019) and ECHAM6.3-HAM2.3-MOZ1.0 with sectional aerosol module SALSA2.0 (hereafter ECHAM-SALSA, Stevens et al., 2013; Kokkola et al., 2018; Tegen et al., 2019). Both GCMs are part of the Aerosol Comparisons between Observations and Models (AeroCom) Phase III GCM Trajectory Experiment (GCMTraj) in which a comparison between the GCMs against reanalysis meteorology was conducted for the years between 2009 and 2013. In this study the simulations for UKESM1 and ECHAM-SALSA cover the years from 2005 to 2018 which are also available

127 from the observations. Comparison between modal (UKESM1) and sectional (ECHAM-SALSA) approaches for
128 estimating the aerosol microphysics provides additional insight into the model behaviour via this Lagrangian evaluation
129 approach. The Hybrid Single-Particle Lagrangian Integrated Trajectory model (HYSPLIT; Draxler and Hess, 1998; Stein
130 et al., 2015) is employed to obtain the backward air mass trajectories. A key difference between our study and previous
131 work, including Isokääntä et al. (2022), is our focus on stratiform precipitation rather than total precipitation. Stratiform
132 precipitation is the dominant type in mid- and high-latitude regions (30–60° from the equator and poleward), whereas
133 tropical regions are typically influenced by convective systems (e.g., Schumacher and Funk, 2023). Since our study area
134 is primarily the boreal forest region of northern Europe, stratiform precipitation is most relevant. The differing impacts
135 of precipitation types on aerosols have also been highlighted by Khadir et al. (2023), who showed that recent tropical
136 precipitation—largely convective—can be linked to downdrafts that transport small particles from higher altitudes to the
137 boundary layer (see also Franco et al., 2022; Machado et al., 2021; McCoy et al., 2021; Williamson et al., 2019).

138 The aim of our research can be summarized into two main objectives (1-2):

- 139 1. Do the relationships between aerosols and experienced precipitation during transport differ between the
140 measurements and GCMs and what are the drivers for the observed differences?
- 141 2. Do the GCMs exhibit similar increase in sulfate mass due to in-cloud production as the observations and are the
142 observed effects reasonable when reflected to model parametrizations?

143 We start out investigation in Sect. 2 by first introducing the observational datasets, followed by summarising the GCM
144 simulations along with details on the air mass trajectory calculations and data co-locations employed in this work. The
145 aerosol properties at SMEAR II are given in Sect. 3 as a necessary background for the following Lagrangian analysis.
146 The relationships between precipitation and aerosol mass and number in the Lagrangian framework are presented first
147 (Sect. 4.1-4.3), followed by a process-chain type evaluation (Sect. 4.4) to understand the driving forces in the
148 relationships. Finally, in Sect. 5, the effects of aqueous-phase processing are presented, followed by overall conclusions
149 (Sect. 6) and outlook (Sect. 7).

150 **2 Data and methods**

151 **2.1 Observations at SMEAR II**

152 Observational data used in this study include long-term measurements of aerosol number size distributions and particle
153 chemistry from SMEAR II (Station for Measuring Ecosystem-Atmosphere Relations in; Hari and Kulmala, 2005) and are
154 described in detail in Isokääntä et al. (2022) and the references therein. SMEAR II station (Hyytiälä, Finland) is classified
155 as a rural environment, surrounded by relatively homogenous Scots pine (*Pinus sylvestris*) forest. In this work particle
156 number size measurements (covering particle diameters between 3–1000 nm) obtained with a differential mobility particle
157 sizer (DMPS, e.g., Aalto et al., 2001) are utilized. Chemical composition (organics, sulfate, and equivalent black carbon)
158 of the particles in the sub-micron range were derived from an aethalometer (e.g., Drinovec et al., 2015) and aerosol
159 chemical speciation monitor (ACSM, Ng et al., 2011). The dataset for particle number size measurements spans 2005–
160 2018, slightly shorter than in Isokääntä et al. (2022), to match the GCM simulation period. The ASCM data extends from
161 2012 to 2018.

162 **2.2 Summaries of the GCMs used in this study**

163 **2.2.1 UKESM1**

164 The United Kingdom Earth System Model (UKESM1) configuration used in this study uses the atmospheric and land
165 components following the protocol set by the Atmospheric Model Intercomparison Project (AMIP, Eyring et al., 2016).
166 The atmospheric component of the model is based on the Global Atmosphere 7.1 (GA7.1) and the Global Land 7.0
167 (GL7.0) configurations, as described by Walters et al. (2019). These are part of the Hadley Centre Global Environment
168 Model version 3 (HadGEM3; Hewitt et al., 2011), which is coupled to the terrestrial carbon/nitrogen cycles (Sellar et al.,
169 2019). It includes interactive stratosphere–troposphere chemistry from the from the UK Chemistry and Aerosol (UKCA)
170 model (Archibald et al., 2020; Morgenstern et al., 2009; O'Connor et al., 2014).

171 Following the AMIP protocol, sea surface temperature and sea ice are taken from the unmodified dataset of Durack et al.
172 (2017) and horizontally interpolated to the model resolution. In this setup, the dynamic vegetation model (Cox, 2001) is
173 turned off. Instead, prescribed vegetation from a historical coupled UKESM1 simulation is used to maintain consistent
174 land-use forcing between the coupled and AMIP experiments. In a similar fashion, seawater concentrations of dimethyl
175 sulfide (DMS) and chlorophyll-a monthly climatologies are taken from the coupled historical experiment and are used by
176 the atmosphere model top calculates fluxes of DMS and primary marine organic aerosol (Mulcahy et al., 2020).

177 The simulations were nudged to ERA-Interim reanalysis (Dee et al., 2011; Telford et al., 2008) u/v (horizontal and
178 vertical), wind fields and surface pressure following the setup design for the AeroCom GCMTraj phase III experiment.
179 The model resolution for these configurations was $1.875^\circ \times 1.25^\circ$ longitude–latitude, corresponding to a horizontal
180 resolution of ~135 km in the midlatitudes. The model has 85 vertical levels which are divided such that 50 levels are
181 between 0 and 18 km and the remaining 35 levels cover heights between 18 and 85 km.

182 Atmospheric composition within UKESM1 is implemented as part of the UKCA model. Within UKCA, the Global Model
183 of Aerosol Processes (GLOMAP; Mann et al., 2010; Mulcahy et al., 2020) is used. This scheme simulates multicomponent
184 global aerosols, including, for example, sulfate, black carbon, and organic matter. The aerosol particle size distribution is
185 represented using five log-normal modes, nucleation soluble, Aitken soluble, accumulation soluble, coarse soluble and
186 Aitken insoluble visualized in Figure S1. More details, including the size ranges for each aerosol mode, are presented in

187 Sect. S1.1. The GLOMAP model also includes various microphysical processes that affect the evolution of aerosol
188 properties. Wet scavenging processes in UKESM1, including below-cloud (impaction), in-cloud (nucleation) and plume
189 scavenging are summarized in Sect. S2 and references therein. As a key difference to ECHAM-SALSA (Sect. 2.2.2)
190 concerning the aerosol parametrizations, new particle formation in the boundary layer is not yet implemented in UKESM1
191 (Mulcahy et al., 2020).

192 For this study the AeroCom GCMTraj UKESM1 simulations (2009-2013) were ran longer to cover years from 2005 to
193 2018 to facilitate robust statistical comparison with the aerosol size distributions and composition measurements obtained
194 from SMEAR II. The model output fields were extracted at high temporal resolution (3-hourly output) for all model levels
195 (when available, otherwise noted as surface). The diagnostics fields utilized in this work (see also Table S4) are aerosol
196 particle size distribution variables (number concentrations and dry diameters for each aerosol mode), chemical
197 components including mass mixing ratios of sulfate noted here as SO_4 (extracted as sulfuric acid H_2SO_4 and then
198 converted, see Sect. S1.1), organic matter (noted here as OA) and black carbon (BC), total (including both liquid rain and
199 snow) stratiform and convective precipitation at the surface, dry air density, sub-grid scale updraught velocity, number
200 of activated particles, total precipitation at the surface, relative humidity and cloud fractions. Additionally, from
201 UKESM1, wet scavenging coefficients (representing removal within the whole atmospheric column) for the different
202 removal processes (nucleation, impaction and plume) and species (OA, H_2SO_4 and BC), SO_2 concentrations, and both
203 vertically resolved and surface liquid stratiform precipitation are inspected. These variables and/or variables derived from
204 them are co-located to the UKESM1 derived HYSPLIT back-trajectories as described in Sect. 2.3.

205 2.2.2 ECHAM-SALSA

206 ECHAM6.3-HAM2.3-MOZ1.0 is a global aerosol-chemistry-climate model consisting of the atmospheric general
207 circulation model ECHAM (Stevens et al., 2013) coupled with the Hamburg Aerosol Model HAM (Tegen et al., 2019)
208 and chemistry model MOZ (Schultz et al., 2018). For this work, as for UKESM1, simulations follow AMIP style runs
209 following the AeroCom phase III GCMTraj experiment setup. Therefore, as for UKESM1, the u/v wind fields and surface
210 pressure were nudged towards ERA-Interim reanalysis data. In addition, the sea surface temperature and sea ice cover
211 were prescribed based on monthly mean climatologies obtained from the AMIP project (Eyring et al., 2016). The model
212 solves atmospheric circulation with vertical gridding of 47 layers extending roughly up to 80 km. Model horizontal
213 resolution for these configurations is $1.875^\circ \times 1.875^\circ$ longitude–latitude.

214 ECHAM6.3-HAM2.3-MOZ1.0 is paired with the sectional aerosol microphysics model SALSA2.0 (ECHAM-SALSA)
215 in which the size distribution is divided into 3 subranges ($d_{p1} = 3 - 50$ nm, $d_{p2} = 50 - 700$ nm and $d_{p3} = 700$ nm – 10 μm)
216 including 10 size classes in logarithmical size space. Subranges d_{p2} and d_{p3} include parallel size classes for insoluble and
217 soluble aerosol species, making the total number of size classes 17 (Kokkola et al., 2018), visualized in Figure S1. More
218 details of the subranges and their compositions are given in Sect. S1.2. Additional details of the aerosol processes
219 calculated in SALSA2.0 can be found in Kokkola et al. (2018) and Holopainen et al. (2020). Wet scavenging
220 parametrizations are summarized in Sect. S2 for below- and in-cloud scavenging.

221 As for UKESM1, simulations cover the years from 2005 to 2018 for ECHAM-SALSA. Data output is also 3-hourly and
222 vertically resolved unless the variable is noted as surface variable. The diagnostics extracted from ECHAM-SALSA (see
223 also Table S4) include aerosol particle size distribution variables (number concentrations and dry diameters for each size
224 class), chemical components including mass mixing ratios of sulfate (SO_4), organics (noted here as OA) and black carbon

(BC), total (including both liquid rain and snow) stratiform and convective precipitation at the surface, dry air density, sub-grid scale updraught velocity, number of activated particles, total precipitation at the surface, relative humidity and cloud fractions. Similar to UKESM1, these variables and/or variables calculated from them are co-located to the ECHAM-SALSA derived HYSPLIT back-trajectories as described in Sect. 2.3.

2.3 Air mass trajectory calculations and data co-location

2.3.1 HYSPLIT

The 4-day (96 h) back trajectories arriving at SMEAR II were calculated by version 5.1.0 of the HYSPLIT (Stein et al., 2015) model for the period from January 2005 to December 2018. The 4-day long back trajectories were used to ensure consistency with the results from Isokääntä et al. (2022). In addition, this is typically a long enough period for slowly moving air masses to travel to the boreal environment from high arctic and marine areas. Arrival height of the trajectories to the receptor station was set to 100 m above the ground level. To obtain the GCM derived trajectories, the meteorological fields from the GCMs were first converted into a consistent netCDF4 format which was then converted into the ARL packed HYSPLIT4 compatible format (Kim et al., 2020). The GCM and ERA-Interim (Dee et al., 2011) reanalysis meteorological datasets required for the HYSPLIT4 trajectory calculations were re-gridded to a consistent 1° horizontal resolution. The vertical discretization of the GCM variables was provided on terrain-following hybrid sigma-pressure levels. In UKESM1, however, the native output is on hybrid height levels, which is not supported by HYSPLIT. Therefore, UKESM1 was output on fixed pressure levels, selected to closely match the ERA-Interim pressure levels.

Trajectories were calculated for every 3rd hour for both reanalysis data and the GCMs, corresponding to GCM output resolution. This led to 8 trajectories per day, a total of 40896 air mass trajectories between 2005-2018 before applying any pre-processing and temporal harmonization of the data (Sect. 2.4). Hereafter, when discussing observational data coupled with the ERA-Interim back-trajectories, those are referred as observations unless mentioned otherwise. It should be noted that reanalysis data is not interchangeable with observations but is used as a proxy in this study.

2.3.2 Co-location of GCM data along the air mass trajectories

The variables from the GCMs described in Sect. 2.2.1 and 2.2.2 were temporally (time), spatially (latitude, longitude) and vertically (variables which covered different model or pressure levels) co-located to the GCM derived air mass trajectories. In short, a co-locator tool (Kim et al., 2020) based off the Community Intercomparison Suite (CIS, Watson-Parris et al., 2016) was used to co-locate 4-dimensional data which uses hybrid altitude coordinates. As the default interpolator within CIS has often difficulties co-locating to the near-surface trajectory points (due to surrounding grid-boxes being at the boundaries of the data domain), the modified co-locator provided more flexibility for the interpolation of these near-surface points. This is relevant also in this work, as for our surface site the trajectories can also travel at low altitudes. In this improved co-locator, when the linear interpolation in the near-surface trajectories would result into a missing value, nearest-neighbour interpolation is used instead. Thus, extrapolation of values can be avoided and information for trajectory points that are within the data domain retained. The co-located GCM data from the air mass trajectory arrival times, i.e., times when the air mass is located at SMEAR II, are used to represent the conditions at SMEAR II, facilitating direct comparison to observational data obtained at the site.

A difference to Isokääntä et al. (2022) where the ERA-Interim precipitation internally processed by HYSPLIT onto trajectories coordinates was used, is that the raw precipitation fields from ERA-Interim are employed in this work by co-

262 locating them to the air mass trajectories in a post-processing step (as for the variables extracted from GCMs described
263 above). This approach was chosen to retain the original numerical precision of ERA-Interim (and GCM) precipitation
264 data, ensuring accurate alignment with co-located GCM variables (e.g., aerosol size distributions and chemical
265 composition), which HYSPLIT does not provide. Here, “consistency” refers to numerical accuracy rather than matching
266 data sources.

267 **2.4 Data harmonization between measurements and GCMs**

268 **2.4.1 Temporal co-location and data pre-processing**

269 The data from the measurements (1-hourly averages) conducted at SMEAR II was temporally co-located with the ERA-
270 Interim derived back-trajectory arrival times (3-hourly). Additionally, the GCM derived trajectories (3-hourly) were only
271 co-located with the times when aerosol observations were available. By adopting this approach, only GCM trajectories
272 corresponding to existing data points in observations were retained and utilized in further analysis. The importance of
273 temporal co-location for model evaluation is discussed, for example, in Schutgens et al. (2016). Harmonisation of the
274 measured aerosol size distribution and composition with the corresponding variables available from the GCMs are
275 described in Sect. 2.4.2 and 2.4.3.

276 For consistency with Isokääntä et al. (2022) identical pre-processing is applied here to the in-situ aerosol observations
277 before the temporal co-location described above. Thus, data points for which the measured wind direction was between
278 120 and 140 degrees were removed due to possible influence of strong VOC (volatile organic compound) emissions from
279 the local sawmill (Heikkinen et al., 2020; Liao et al., 2011). In addition, trajectories crossing the area of Kola Peninsula
280 were excluded as in Isokääntä et al., (2022) due to strong pollution sources within the area (Heikkinen et al., 2020;
281 Kulmala et al., 2000; Riuttanen et al., 2013). This led to aerosol size distribution data covering the years between 2005
282 and 2018 (number of final data rows/trajectories: 30688) and aerosol chemical composition for the years between 2012
283 and 2018 (number of final data rows/trajectories: 6174). Distribution of the data points over the years are shown in Figures
284 S2 and S3.

285 **2.4.2 Aerosol particle number size distribution**

286 The DMPS (differential mobility particle sizer, e.g., Aalto et al., 2001) observations include 51 size bins in the observed
287 size range ($d_p = 3\text{-}1000$ nm). For UKESM1, complete log-normal particle number size distributions (Seinfeld and Pandis,
288 2016) were calculated by using the modal parameters (dry diameters, number concentrations and geometric mean
289 diameters) given by the model. The number size distribution is discretised into the same size grid as the observations i.e.,
290 the bin midpoints are identical to the ones available from the DMPS measurements. This approach was possible as in
291 SMEAR II the size grid DMPS applies stays constant over the whole investigated period. This harmonization was
292 conducted for each hour along the air mass trajectories using the co-location approach described in Sect. 2.3.2 as
293 UKESM1 provided all needed modal parameters for calculation of the full particle number size distributions (PNSD)
294 along the trajectories.

295 For ECHAM-SALSA, the number concentrations of soluble and insoluble bins (i.e., size classes) were added together for
296 each size bin. To make the logarithmic number size distribution comparable to UKESM1 data and DMPS measurements,
297 the values within each size bin (i) were divided by the logarithm of the maximum size $d_{i,\max}$ minus the logarithm of the
298 minimum size $d_{i,\min}$ i.e., by $\log_{10}(d_{i,\max}) - \log_{10}(d_{i,\min})$ for that size bin (see Table S3). Similar to UKESM1, this was

299 conducted along the trajectories. For aerosols, ECHAM-SALSA bins ranging from 3.0 nm to 1700 nm in diameter are
300 studied, as by strictly limiting to sub-micron bins (≤ 700 nm), the largest sub-micron particles ($700 \text{ nm} < d_p \leq 1000$ nm)
301 that do contribute to the total particle mass, would be lost. Sensitivity analysis was conducted including only the sub-
302 micron bins, and none of the conclusions changed.

303 Integrated variables, such as total number and mass concentrations (for submicron particles) were calculated from the
304 particle number size distributions by assuming the particles are spherical and have a constant density of $\rho = 1.6 \text{ g cm}^{-3}$.
305 This density corresponds to the average density of particles observed at SMEAR II (e.g., Häkkinen et al., 2012). Again,
306 these quantities were calculated for each hour (i.e., 96 data points, see Sect. 2.3.1) along every single air mass trajectory.

307 2.4.3 Chemical composition

308 Observational data for organic aerosol (hereafter OA) and sulfate (hereafter SO_4) was obtained using observations from
309 ACSM (aerosol chemical speciation monitor, Ng et al., 2011) which is most efficient at measuring particles with ~ 75 -
310 650 nm of vacuum aerodynamic diameter, passing through particles up to $1 \mu\text{m}$ (Liu et al., 2007). For UKESM1, Aitken
311 and accumulation mode are used in this context by summing the mass mixing ratios (MMR, kg of species per kg of air)
312 of these modes, including both soluble and insoluble modes when available. Due to the definition of the modes in
313 UKESM1, these correspond to particle diameters between 10-500 nm (see Sect. S1.1), thus having large overlap with the
314 size range most efficiently represented in ACSM. The MMRs from UKESM1 and ECHAM-SALSA are converted into
315 mass concentrations by multiplying the MMRs with the density of the air to facilitate comparisons to chemistry
316 observations given in the units of $\mu\text{g m}^{-3}$. Equivalent black carbon (hereafter BC) was measured with an aethalometer
317 using a cut off diameter of $10 \mu\text{m}$ (PM_{10}). Due to most of the absorbing particles at SMEAR II being at sub-micron range,
318 the difference in the BC mass between PM_1 and PM_{10} is only 10 % (Luoma et al., 2019). Therefore, from UKESM1,
319 Aitken and accumulation modes are also used to estimate the total BC. In addition, to obtain SO_4 from H_2SO_4 (sulfuric
320 acid) which is the UKESM1 native output, a conversion factor is used (see Sect. S1.1). From ECHAM-SALSA, bins with
321 diameters ranging from 19.6 nm to 700 nm (see Sect. S1.2) are used to estimate the total sub-micron OA, SO_4 and BC,
322 including again both soluble and insoluble bins. Here, for ECHAM-SALSA, the largest bin of which a portion also
323 consists of aerosols larger than $1 \mu\text{m}$ ($700 \text{ nm} < d_p < 1700$ nm) is not included to ensure consistency with the ACSM
324 measuring efficiency (which decreases from ~ 650 nm up to the maximum size of $1 \mu\text{m}$).

3 Aerosol properties at SMEAR II – Eulerian comparison between observations and GCMs

To set the scene and provide context to GCM development since these previous studies (see also e.g., Reddington et al., 2016), a short assessment of the differences and similarities in Eulerian framework between the aerosol observations, UKESM1 and ECHAM-SALSA at SMEAR II is given here. Air mass transport between ERA-Interim and the GCMs is first assessed (Sect. 3.1), followed by the aerosol particle number size distributions (Sect. 3.2) and chemical composition (3.3). This provides the necessary background information to facilitate further comparisons within the Lagrangian evaluation framework used in this work.

3.1 Comparison of air mass transport between ERA-Interim and the GCMs

To ensure the differences shown in the following sections for the Eulerian analysis are not driven by diverging transport pathways between the GCMs and ERA-Interim, the air mass transport routes were inspected. The air mass transport routes in Figure 1 show very similar patterns for ERA-Interim and the GCMs, i.e., the differences are, on average, very small—as expected for simulations in which wind fields are consistently nudged to ERA-Interim reanalysis. Vertical transport differences exist (Figure S5), which can be attributed to potential temperature not being nudged, which follows standard practices (Zhang et al., 2014). For this station, however, these differences are relatively small, and the largest differences are in areas with low frequency of trajectories. Therefore, any observed differences in the analyses presented in the following sections are unlikely to be dominated by differences in the air mass transport.

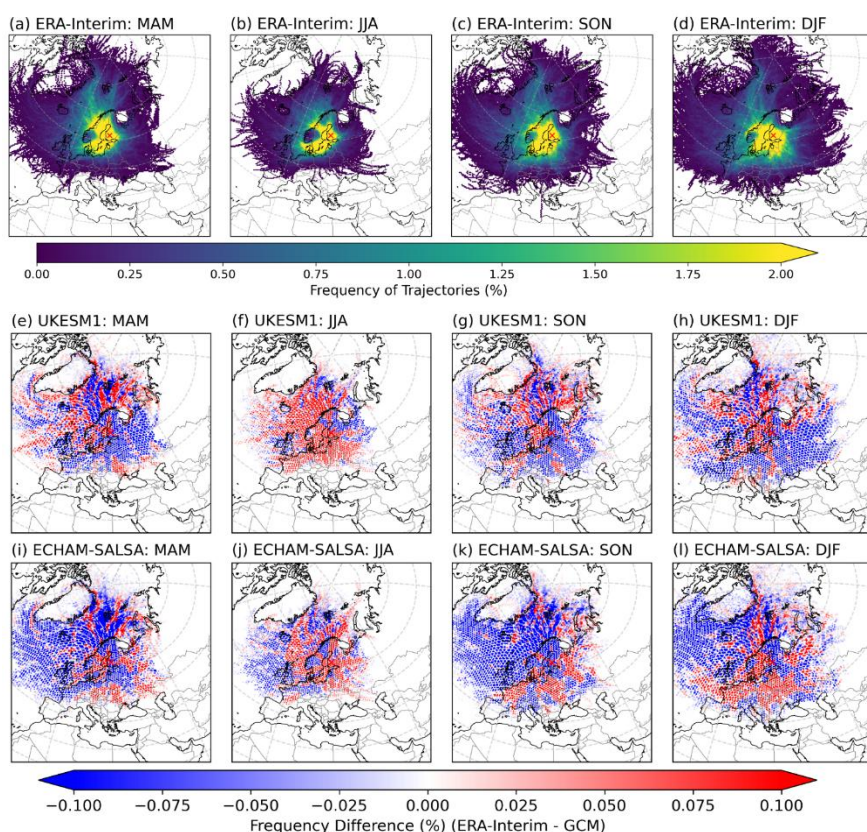


Figure 1 ERA-Interim air mass trajectory frequencies for spring (MAM), summer (JJA), autumn (SON) and winter (DJF) are shown in the top row. Frequencies for UKESM1 (e-h) and ECHAM-SALSA (i-l) are shown as differences to the ERA-Interim. Before calculating the differences, the GCM hexagonal grid (150 hexagons in the x-direction) were first regridded to match the gridding in ERA-Interim. Red cross shows the location of SMEAR II.

3.3 Aerosol particle number size distributions

In Figure 2 particle number size distributions from the GCMs are compared with observational data at SMEAR II. The figure reveals that UKESM1 underestimates the number concentration of the small ($d_p < 50$ nm) particles, especially during summer (Figure 2b, Table S5). This is, however, expected, as the new particle formation from boundary layer nucleation was not implemented in UKESM1 (Mulcahy et al., 2020). ECHAM-SALSA does have a better representation of the PNSD of the smaller aerosol particles during spring and summer when compared to observations (Figure 2c), During warmer seasons, also the absolute number concentrations agree well between observations and ECHAM-SALSA (see nucleation mode from Table S5). This highlights the importance of NPF from nucleation in the boundary layer, especially in summer. During winter, however, ECHAM-SALSA does exhibit some overestimation for Aitken mode aerosols (Figure 2e and Aitken mode from Table S5).

During winter, UKESM1 overestimates larger Aitken and accumulation mode aerosols (d_p up to 200 nm) compared to the observations (Figure 2b and g), but during spring the number concentration of the accumulation mode aerosols is very close to observations (367 cm^{-3} in UKESM1 vs 352 cm^{-3} in observations as shown in Table S5). This is somewhat surprising considering the missing growth of small particles from NPF into accumulation mode, however, this could indicate that there are other processes that dominate the accumulation mode. During winter (Figure 2g) the observations exhibit clear bimodal PNSD peaking around 50 and 200 nm but neither of the GCMs is able to capture this behaviour. Overall, both GCMs tend to be shifted towards the larger sizes in all seasons (Figure 2d-g), and this effect is slightly more pronounced in UKESM1. Overall, ECHAM-SALSA better estimates of the peak values of the PNSD, except in winter (Figure 1g), when it overestimates the particle concentrations at the size range of $d_p = 50 - 100$ nm.

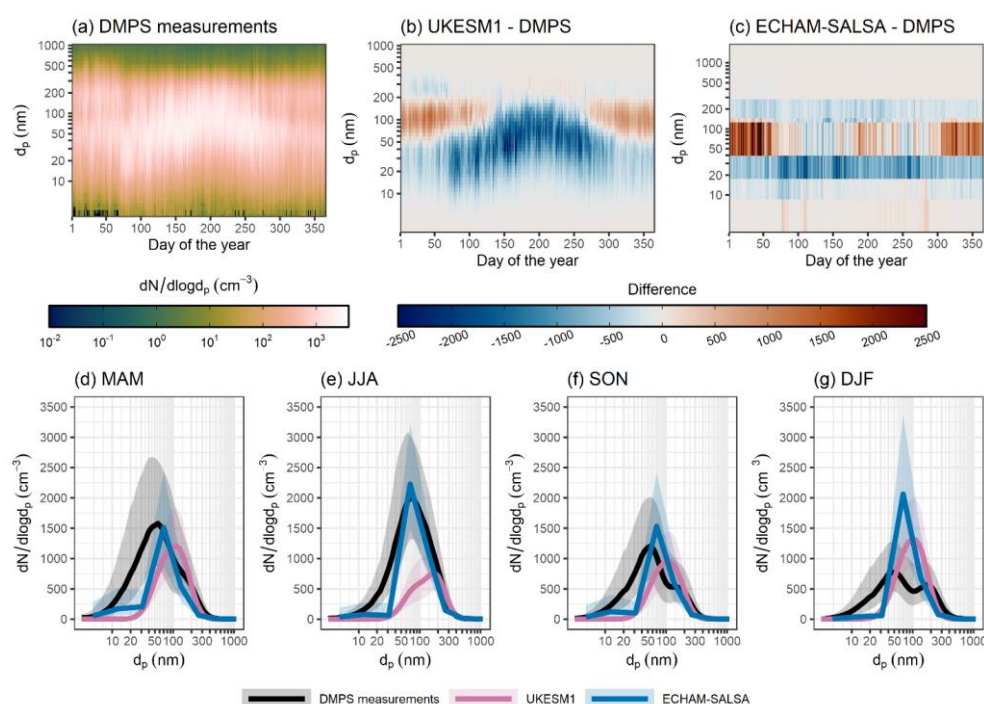


Figure 2 Particle number size distribution at SMEAR II as medians for the day of the year for DMPS measurements (ground level) are shown in (a), followed by the differences between the DMPS observations and the GCMs in (b) and (c). For subplot (c), the measured size distribution was first regridded to the ECHAM-SALSA bins by integrating between the upper and lower limit of each ECHAM-SALSA size bins before calculating the difference. Median PNSDs for each season are shown in (d)-(g) with shaded areas indicating the 25th and 75th percentiles.

3.3 Chemical composition of the aerosols

Particle chemical composition as a mass concentration for each chemical species from the composition measurements and the GCMs at SMEAR II (trajectory receptor location) is illustrated in Figure 3, and the seasonal patterns are typical for this location. Largest concentration of organic material is present during summer (JJA) and smallest in winter (DJF). Both GCMs also have pronounced OA concentration during summer compared to the other seasons, and UKESM1 captures the pronounced OA concentrations observed during summer particularly well (median OA $2.0 \mu\text{g m}^{-3}$ and $2.2 \mu\text{g m}^{-3}$ in UKESM1 and observations, respectively, Table S6). A portion of the small underestimation of the OA concentrations of the GCMs during spring and summer could, however, be influenced by the height of the observations as chemical composition measurements are conducted at the surface whereas the GCM data shown here are at the trajectory arrival point height at the receptor station (100 m.a.g.l.). Scale difference likely also plays a role, as the point measurements are compared with the GCM grid box values interpolated to air mass trajectories. Monthly data (Figure 3e) shows the second OA peak for the observations to be in February, as expected based on Heikkinen et al. (2020), and in ECHAM-SALSA this peak falls on January. UKESM1 peaks in February, but the difference in the concentrations (compared to observations) between February and January/March is very small. The seasonality of the OA concentrations presented here for both observations and GCMs also agrees with the results from Blichner et al. (2024) who presented the same GCMs but for a different time period. Differences in the monthly peak concentration can be observed for BC too, where observations and UKESM1 peak in February, but ECHAM-SALSA exhibits the largest BC concentrations in January (Figure 3g).

In general, even though a perfect harmonization of the particle chemical composition data between observations and GCMs is not achieved (see Sect. 2.4.3), the median concentrations between observations and GCMs agree relatively well when the overall seasonality is inspected (Figure 3a-d); the concentrations are dominated by OA in all seasons, followed by SO_4 and BC. Inspection of the monthly median concentrations (Figure 3e-g), however, revealed that differences also exist.

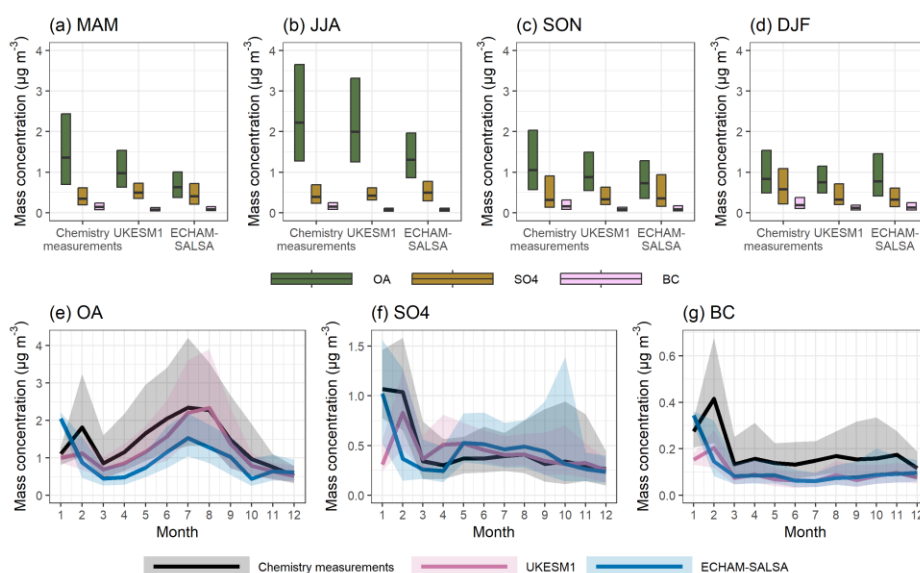


Figure 3 Average seasonal mass concentration of sub-micron OA, SO_4 and BC at SMEAR II from the chemical composition measurements, UKESM1 and ECHAM-SALSA is shown in (a)-(d). Black horizontal lines show the median and the boxes extend between 25th and 75th percentiles. Monthly median (lines) concentrations and 25th-75th percentiles (shaded areas) are presented in (e)-(g).

399 **4 Lagrangian analysis of overall effects of integral precipitation on aerosols at SMEAR II**

400 In this section we use the Lagrangian framework to investigate the potential wet removal of the aerosols. In Sect. 4.1 we
401 first examine the impact of using vertically resolved liquid precipitation (UKESM1 only), which has not previously been
402 done for Lagrangian trajectory analyses. Then we inspect the relationship between accumulated precipitation and aerosols
403 for the two GCM s used in this study: UKESM 1 and ECHAM-SALSA. In Sect. 4.2 we focus on total aerosol mass and
404 number, and in Sect. 4.3 we focus on the OA, BC, and SO₄ portions of the total mass for submicron-size aerosols. Then,
405 in Sect. 4.4, the processes controlling the precipitation-aerosol relationships presented in the previous sections are
406 investigated, and the differences are discussed in detail *between the GCMs* (Sect. 4.4.1) and *within each GCM* (Sect.
407 4.4.2). Supplementary analysis assesses the representability of the models employed here amongst larger group of GCMs
408 (Sect. S4).

409 **4.1 Assessment of surface vs. vertically resolved precipitation in Lagrangian wet removal**

410 In earlier studies assessing aerosol-precipitation relationships at SMEAR II using the Lagrangian framework (e.g.,
411 Isokääntä et al., 2022; Khadir et al., 2023; Tunved et al., 2013) the vertical position of the trajectories with respect to the
412 precipitating clouds was not considered. The approach, therefore, does not allow for separation between in-cloud and
413 below-cloud precipitation scavenging. Instead, it provides us with the overall effect of precipitation (hereafter noted as
414 wet removal), in which the surface precipitation is used as a proxy for the experienced precipitation by the air mass. This
415 also means that it could include trajectories that travel above the precipitation, potentially confounding interpretation of
416 the results.

417 For this study, the impact of this simplification was examined by extracting the vertically resolved liquid precipitation
418 from UKESM1, which can be compared to the surface precipitation (see Appendix A). Based on this analysis, it was
419 possible to conclude (see e.g., Figure A1) that for this station the surface precipitation is a relatively good proxy for the
420 experienced precipitation by the air mass. Therefore, and to be able to include the effects due to snowfall, which was
421 unfortunately not extracted with high enough vertical resolution from UKESM1, the surface precipitation is continued to
422 be used in this study. Vertically resolved precipitation was not available from ECHAM-SALSA.

423 **4.2 Relationship between precipitation and aerosol mass and number concentrations**

424 .The removal of the normalized masses ($dp = 3-1000$ nm, Figure 4a) by accumulated stratiform precipitation for
425 observations and both GCMs exhibit exponential decrease reaching asymptotic behaviour after ~ 10 mm of accumulated
426 precipitation (after 5 mm for UKESM1 during summer). Normalization of the median mass/number concentration to the
427 median value under zero accumulated stratiform precipitation is used in this study. This approach aims to minimize the
428 influence of differences in the native particle number size distributions (e.g., Figure 1), which affect total mass and
429 number concentrations, and instead highlight the removal attributable to precipitation.

430 For the particle number concentration ($dp = 3-1000$ nm), there are clear seasonal differences (Figure 4b). ECHAM-
431 SALSA and the observations show clear seasonal differences in particle number removal, with much more efficient
432 removal in winter than in summer. UKESM1, however, does not display this seasonal contrast—likely because it lacks
433 boundary layer nucleation, a key source of small particles during summer, which leads to similar particle number
434 concentrations across seasons. Inspection of the seasonality is relevant, as differences in the relationships could be driven
435 by different particle size distributions at the station which vary by season due to differences in meteorology (e.g., origin

of air-masses, temperature and sunlight) along the air mass trajectories. Seasonality also impacts to the type of the precipitation (liquid vs snow and stratiform vs convective, for example).

Figure 4c shows that the seasonal patterns (e.g., more samples for smaller precipitation values in summer) in the distribution of accumulated precipitation are similar for both models and observations, thus unlikely driving differences in the aerosol-precipitation relationships. The relationships between the aerosol mass, number, and mean stratiform rainfall rate along the trajectory (Figure S6a-b) exhibit similar seasonal differences as the relationships in Figure 4a-b. For example, in summer, UKESM1 exhibits the strongest initial reduction for particle mass (Figure S7a). Observations and ECHAM-SALSA exhibit minimal to no reduction or particle number during summer (Figure S7b), similar to Figure 4b. Non-normalized mass and number concentrations are shown in Figures S7 and S8.

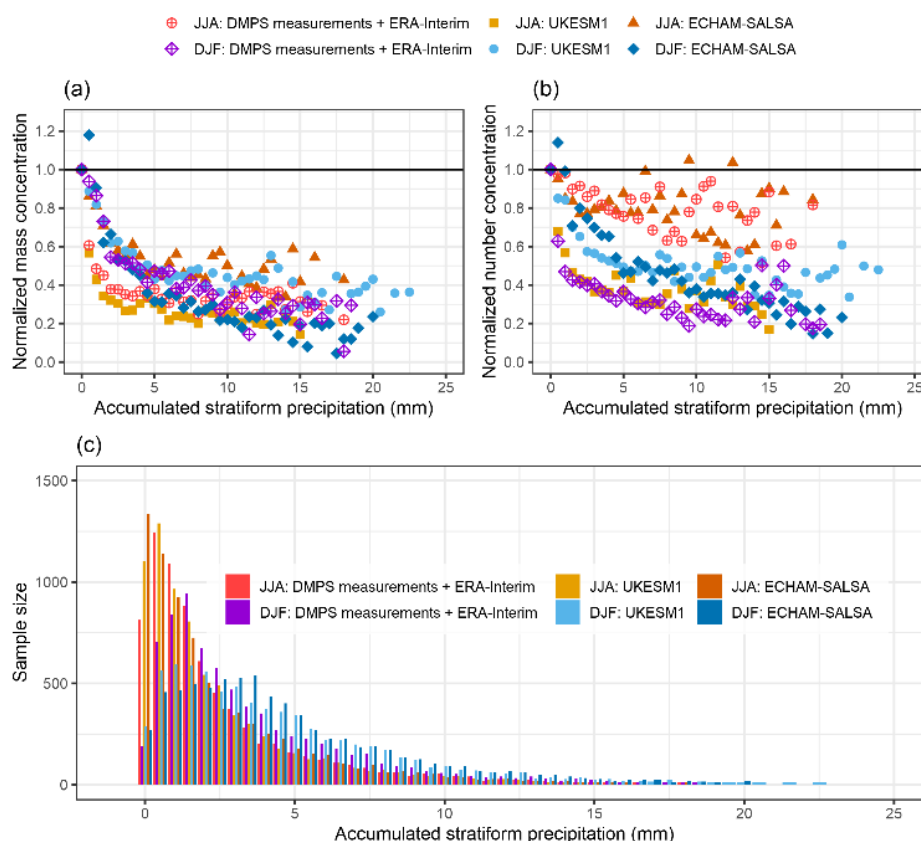


Figure 4 Normalized total ($d_p = 3-1000$ nm) particle mass (a) and number (b) at SMEAR II for summer (June, July and August) and wintertime (December, January and February) as a function of accumulated stratiform surface precipitation (incl. both liquid and snow) along the 96 hour long air mass trajectories for observations (DMPS measurements paired with ERA-Interim trajectories) and GCMs. The coloured points show the median values for each 0.5 mm bin of accumulated precipitation when the number of trajectories in the bin was 10 or larger. The sample size for each corresponding bin is shown in (c).

4.3 Relationship between precipitation and aerosol chemical composition

The normalized masses of OA, BC, and SO_4 in submicron-sized particles as a function of accumulated stratiform precipitation (including both liquid and snow) for the observations and the GCMs is shown in Figure 5 (see also Figure S9 showing the same data but grouped differently for easier comparison between the species). The division into warmer and colder months follows the monthly median temperatures (measured at the site) as in Isokääntä et al. (2022). The

sample sizes in Figure 5g-h agree well during warmer months between the GCMs. During colder months (Figure 5h) more differences emerge for the smaller precipitation bins (< 3 mm of accumulated precipitation).

The general patterns between the observations and GCMs are similar for all species—exponential decrease is observed for the mass concentrations, similar to the relationships between total particle mass and precipitation shown in Figure 4a. The seasonal differences for the total particle mass (Figure 4a) and the chemical constituents are comparable despite the different approach used to separate the data into temperature regimes instead of seasons. During the colder months (Figure 5d-f), ECHAM-SALSA exhibits the most efficient reduction for all the three species, as expected based on the reduction of the total aerosol mass (Figure 4a). During the warmer months (Figure 5-c), UKESM1 tends to show more efficient reduction than ECHAM-SALSA, the effect being most pronounced for OA. This is in line with the derived reduction of total particle mass and number during summer shown in Sect. 4.1 (Figure 4a-b), in which ECHAM-SALSA exhibited stronger reduction during winter and UKESM1 during the summer.

The observational data presented by Isokääntä et al. (2022) showed that the reduction of SO_4 due to accumulated total precipitation in the warmer months was less efficient compared to other species, despite SO_4 being highly hygroscopic and thus relatively easily activated as a cloud droplet. This is relevant also in this study, as the activation into cloud droplets followed by precipitation is the dominant reduction mechanisms also for the mass of the different chemical species (discussed in more detail in Sect. 4.4). Similar to Isokääntä et al. (2022), the derived reduction for SO_4 is less efficient (i.e., smaller end concentrations are reached) compared to OA and BC also here for the observations and UKESM1 (Figure S9a-b), though the differences between species are overall smaller but still statistically significant (Kruskal-Wallis rank sum test, $p < 0.001$). For ECHAM-SALSA, the derived removals between OA and SO_4 do not differ (Figure S9c, Kruskal-Wallis rank sum test, $p = 0.2$) during warmer months, but BC shows more efficient reduction with the accumulated stratiform precipitation than OA and SO_4 . This could be arising from the fact that, in ECHAM-SALSA, all BC is basically in the soluble particles (Figure S10b) but OA and SO_4 can reside in the insoluble particles as well.

Isokääntä et al. (2022) hypothesized that the low derived removal efficiency of SO_4 during warmer months could be caused by the species being distributed to different sizes depending on the season. Inspection of the size resolved chemical composition from the GCMs (Figure S10), however, is not able to fully explain the observed seasonal differences: SO_4 in the GCMs is almost completely distributed to the soluble accumulation mode, and the seasonal differences are only minor. In ECHAM-SALSA, small contribution of insoluble SO_4 in the accumulation mode is present, but the difference between the seasons is small (Figure S10b). Other possible explanations could include, for example (but not limited to), mixing state (internal/external) of the particles and production of SO_4 through cloud processing, which could compensate for the reduction by stratiform precipitation.

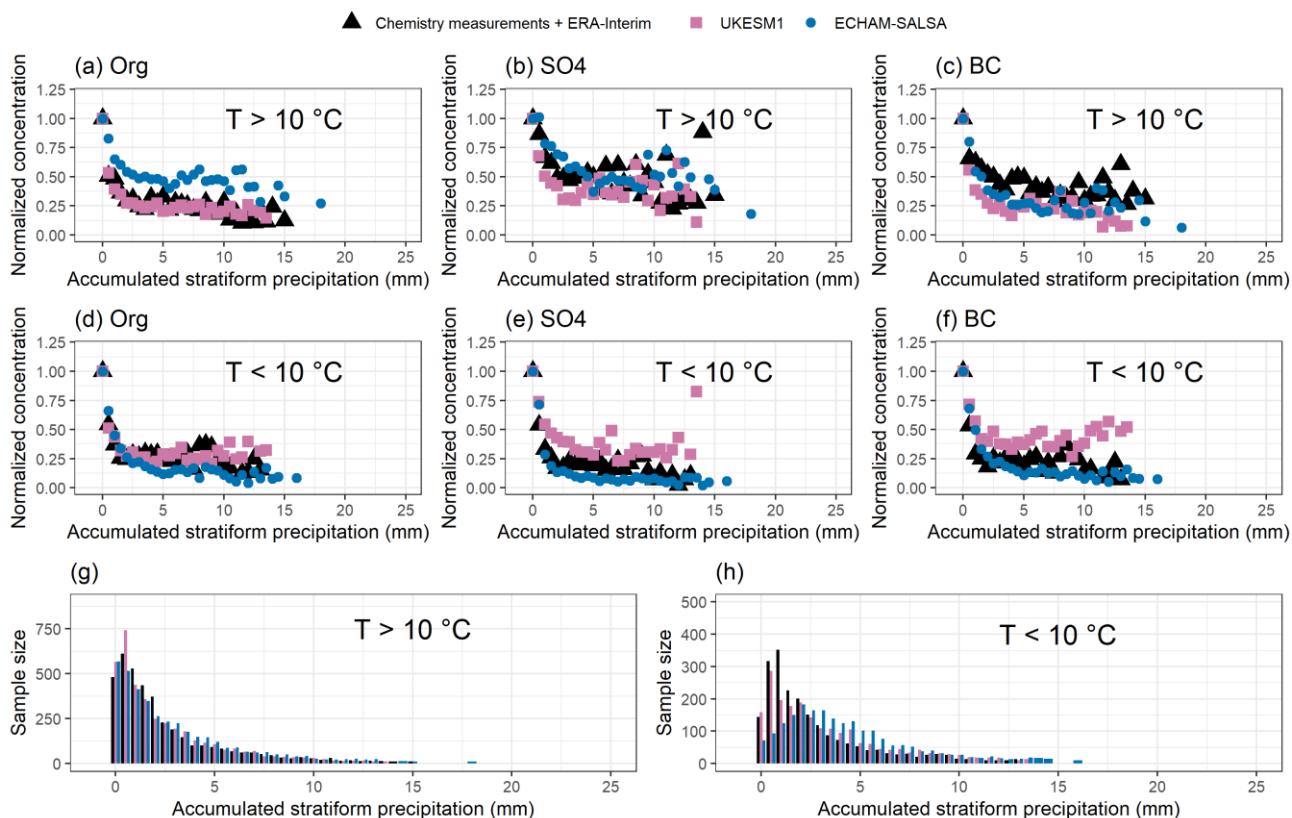


Figure 5 Normalized mass concentration for submicron OA, SO₄ and BC at SMEAR II as a function of accumulated stratiform surface precipitation along the 96 hour long air mass trajectories for observations (chemistry measurements paired with ERA-Interim trajectories) and the GCMs for warm ($T > 10\text{ }^{\circ}\text{C}$, (a)-(c)) and cold ($T < 10\text{ }^{\circ}\text{C}$, (d)-(f)) months. The coloured points show the normalized median values for each 0.5 mm bin of accumulated precipitation when the number of trajectories for the bin was 10 or larger. The sample size for each corresponding 0.5 mm bin is shown in (g)-(h).

4.4 Process-chain evaluation for understanding the relationship between precipitation and aerosols

To understand the differences between GCMs and observations in Figure 4 and Figure 5, we assess the relative importance of wet removal pathways. Prior studies (Isokääntä et al., 2022; Tunved et al., 2013; Wang et al., 2021), suggests that in-cloud scavenging, particle activation followed by rainout, is the dominant removal mechanism for submicron particles in this region. For UKESM1 the relative contributions of the removal types (below-cloud impaction, nucleation followed by rainout, and plume scavenging) were quantified using median scavenging coefficients along the trajectories (see Sect. S2). These scavenging coefficients represent the removal within the total atmospheric column, median values along complete trajectories being 0.040 (JJA) and 0.028 (DJF) moles s^{-1} for impaction, 0.700 (JJA) and 0.191 (DJF) moles s^{-1} for nucleation followed by rainout and 0.001 (JJA) and 0.000 (DJF) moles s^{-1} for plume scavenging.

As shown in Figure 6 for organic aerosol (OA), which dominates the particle mass in SMEAR II, e.g., Heikkinen et al., (2020), nucleation followed by rainout dominates removal. Similar patterns are seen for SO₄ (H₂SO₄) and BC (Figure S11), supporting that in-cloud removal is the main process in this region, consistent with Isokääntä et al. (2022).

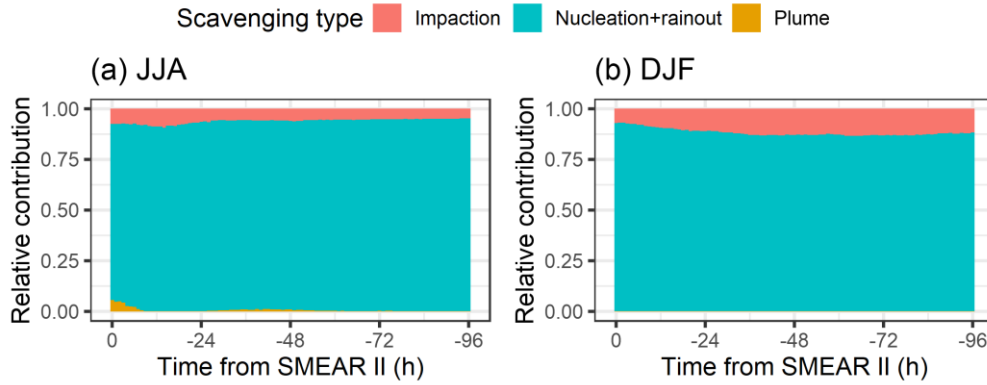


Figure 6 Relative contributions of the different removal pathways in UKESM1 for OA in (a) summer/JJA and (b) winter as a function of time from SMEAR II. Impaction refers to the below-cloud impaction scavenging, nucleation + rainout describes the activation process followed by removal of the particles via the formed raindrops, and plume scavenging is the removal due to convective clouds.

As noted above, nucleation followed by precipitation-driven removal explains the patterns in Figure 4 and 4. To understand differences in this process across models, we compare key variables along air mass trajectories related to in-cloud removal. Previous studies (Dusek et al., 2006; Ohata et al., 2016; Partridge et al., 2012; Reutter et al., 2009) have emphasized the role of sub-grid processes and variables influencing droplet activation, such as particle size and vertical air motion. We therefore examine how model representations of activation—affected by sub-grid vertical velocities and aerosol size distributions—influence removal.

Key variables controlling the aerosol activation into cloud droplets (presented in Figure 7a-j shows the number of particles with diameter > 80 nm (N_{80}) and sub-grid scale vertical velocities (referred as updraught velocities), which control droplet formation. The accumulation mode particles are likely to activate to cloud droplets (Croft et al., 2010; Partridge et al., 2012), and updraught velocities drive supersaturation needed for activation. The activated fraction (N_{act}/N_{tot}) is shown in Figure 7k-o, and the rainfall rates (at the surface) are presented in Figure S12. In addition, total number (N_{tot}) and total mass of the particles (M_{tot}) at the submicron range, a air mass heights and number of activated particles (N_{act}) are presented in Figure S13. Chemical composition, relevant for hygroscopicity and droplet formation, is shown in Figure S15.

Together, these factors determine whether the regime is the aerosol- or updraught limited (Reutter et al., 2009). Figure 4 and Figure 5 showed strong seasonal contrasts, and seasonal differences in N_{80} , updraughts, and activation are also evident during transport (Figure 7). Section 4.3.1 discusses seasonal characteristics within each GCM, followed by a model–observation comparison in Sect. 4.3.2.

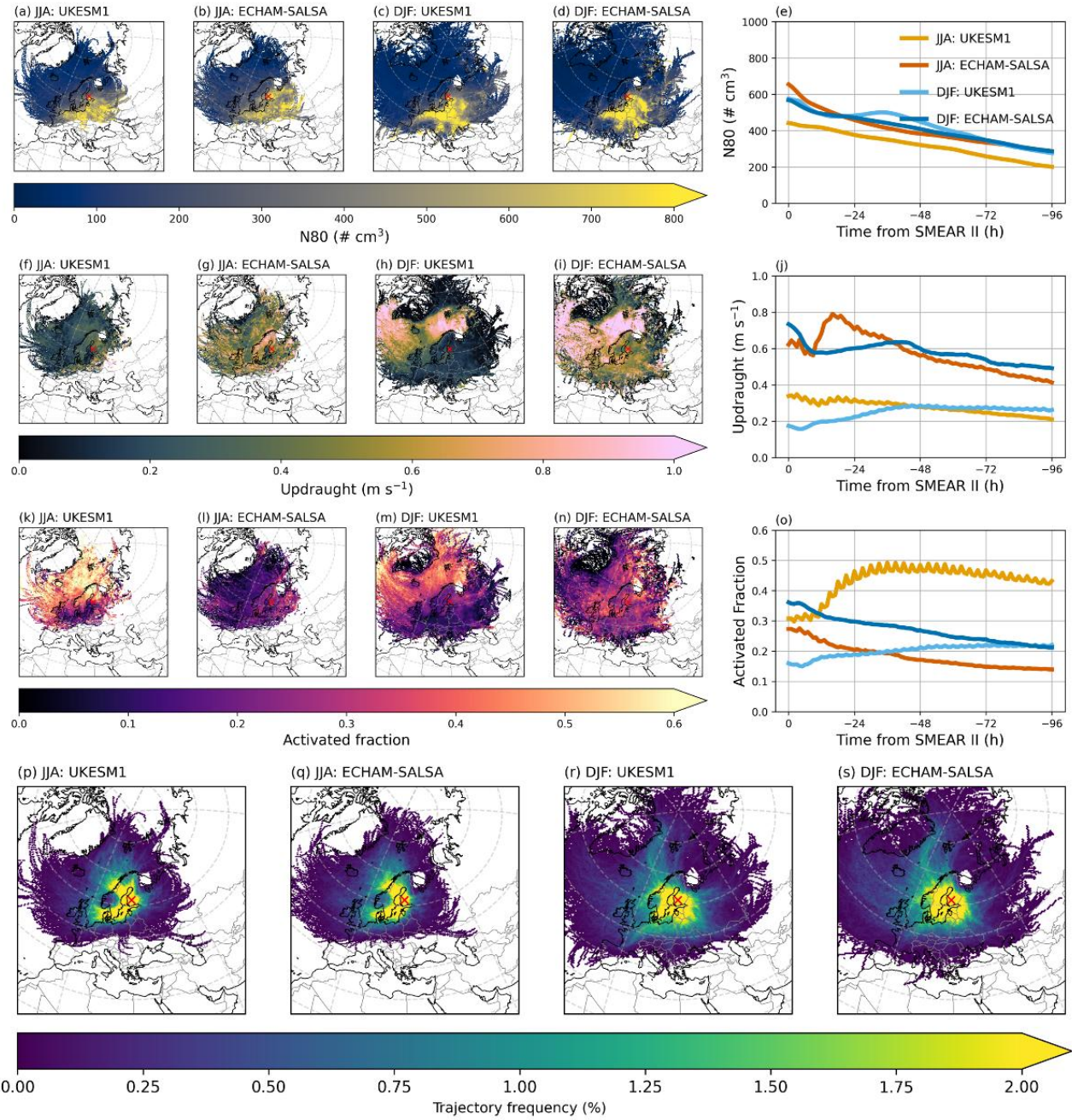


Figure 7 The evolution of the main drivers for the wet removal (nucleation followed by rainout) along the trajectories. The first row from the top displays the N_{80} (number of particles for which $d_p > 80 \text{ nm}$), the second row shows the sub-grid scale updraught velocities (m s^{-1}), third row displays the activated fraction of particles, and the bottom row shows the corresponding trajectory frequencies. For the maps, means are calculated for each hexagonal gridbox (grid resolution being 150 in the x-direction) that the trajectory crosses, and for the rightmost panels, means have been calculated for each hour along the trajectory. For the updraught velocities and activated fractions, only values when trajectory is in-cloud are shown.

4.4.1 Seasonal differences within each GCM

In UKESM1, the derived removal for the particle mass during summer is clearly stronger, especially up to $\sim 10 \text{ mm}$ of accumulated precipitation, compared to winter (Figure 4a). For the particle number, the differences between summer and winter are less pronounced, and similar concentrations at the receptor station are reached (Figure 4b) with high

537 accumulated precipitation. A seasonal difference in the absolute values of N_{80} can be observed, the number concentration
 538 being approximately $100 \text{ particles cm}^{-3}$ larger during winter compared to summer (Figure 7e). This difference, wintertime
 539 values being larger, is also seen in N_{tot} (Figure S13e). As stated in Sect. 2.2.1, the boundary layer nucleation is absent in
 540 UKESM1—a process being especially frequent around SMEAR II during spring and summer (Nieminen et al., 2014).
 541 This is likely the cause for the observed differences in N_{tot} as the model lacks large portion of the smaller particles during
 542 summer. For the mass, however, the summertime M_{tot} is larger (Figure S13j). This could imply that UKESM1 has more
 543 numerous medium-sized particles during summer, or, that on average, the particles in summer are larger compared to
 544 winter, thus having larger contribution to particle mass. Figure 2 supports the latter scenario, showing the average PNSD
 545 at SMEAR II peaking at larger particle sizes in summer compared ($\sim 200 \text{ nm}$, Figure 2g) to winter ($\sim 100 \text{ nm}$, Figure 2i).

546 The seasonal differences between the updraught velocities in UKESM1 are small, until about 48 hours before arrival
 547 (Figure 7j). After that, the summertime updraught velocities exhibit little to no change, but wintertime updraught
 548 velocities decrease as the air mass travels closer to SMEAR II. These differences relatively close to the receptor station
 549 can be attributed to the geographical distribution of the updraught velocities: close to SMEAR II (across Finland, Sweden
 550 and Norway, for example), the values are larger in summertime (Figure 7f) compared to wintertime (Figure 7h). These
 551 regions coincide with areas of high trajectory frequency, meaning most air masses pass through them. As a result, the
 552 elevated updraught velocities in these regions strongly influence the averages shown in Figure 7j. Activated fractions
 553 differ markedly between seasons (Figure 7o), with nearly half of aerosols activating in summer compared to about one
 554 fifth in winter. These seasonal differences align with the spatial patterns of activated fractions and trajectory frequencies
 555 (Figure 7k, p), showing particularly high values over northern Norway and extending into the Arctic Ocean. During
 556 winter, the activated fractions in this area are much lower (Figure 7m). The N_{act} , on the other hand, displays minor
 557 differences between the seasons in UKESM1 but is slightly larger in winter. However, considering the fact that N_{tot} in
 558 UKESM1 is much higher in winter (Figure S13e) as mentioned earlier, the larger activated fraction (derived as $N_{\text{act}}/N_{\text{tot}}$)
 559 in summer is reasonable.

560 The chemical composition of particles during their travel in UKESM1 (Figure S14a) reveals that overall, during summer,
 561 the mass concentration is completely dominated by soluble modes, whereas in winter, a portion of insoluble OA in the
 562 Aitken mode is also present. Soluble SO_4 in the accumulation mode contributes more in winter, but this is greatly
 563 compensated by soluble OA in both Aitken and accumulation modes during summer. If the higher solubility of OA in
 564 summer compensates for the lower SO_4 levels, this could further enhance the particle activation potential in UKESM1
 565 during summer compared to winter. Figure 8 shows the relationship between mean activated fraction and mean updraught
 566 velocity that the air mass experienced before arriving at SMEAR II for the summer and winter. For UKESM1, the
 567 relationship between these two variables is clearly stronger in summer (slope of 2.12, Figure 8a) compared to winter
 568 (slope 0.62, Figure 8b). Therefore, during summer, even a very small increase in updraught could cause a very large
 569 increase in the activated fraction. Due to this, the slightly higher updraught velocities during summer, when the air masses
 570 approach SMEAR II (Figure 7j), could play a major role, eventually also leading to the larger activated fractions during
 571 summer. This, together with the points discussed above (such as the availability of CCN, N_{tot} and particle chemistry along
 572 the trajectories), likely causes the seasonal differences observed in the reduction of particle mass in Figure 4a. When also
 573 considering the missing boundary layer nucleation in UKESM1 as mentioned earlier, lack of seasonality in the derived
 574 removal of total particle number in UKESM1 (Figure 4b) can also be explained.

575 **ECHAM-SALSA** exhibits stronger reduction (i.e., lower concentrations are reached with increasing accumulated
 576 precipitation) during winter than in summer for both particle mass (Figure 4a) and number (Figure 4b). The number of
 577 particles for which $80 \text{ nm} < d_p \leq 1000 \text{ nm}$ (N_{80}) is relatively similar between summer and winter, exhibiting increase from
 578 ~ 300 up to $\sim 650 \text{ particles cm}^{-3}$ as the air mass reaches SMEAR II. During summer, the N_{tot} in ECHAM-SALSA is clearly
 579 larger compared to winter (Figure S13e). This is expected due to the strong contribution of small aerosols during summer
 580 (e.g., Figure 2c). The total mass (M_{tot}), however, is relatively alike between the seasons (Figure S13j), which is reasonable
 581 due to the similar contribution of N_{80} in both seasons, as these particles mostly contribute to particle mass.

582 The updraught velocities in ECHAM-SALSA exhibit large location-dependent seasonal differences (Figure 7g versus i),
 583 especially over the oceans, where the updraught velocities are larger during winter (Figure 7i) than in summer (Figure
 584 7g). However, overall, the average experienced updraught velocities during the transport are rather similar in magnitude
 585 between the two seasons (Figure 7j). This overall similarity occurs because the frequency of trajectories passing over the
 586 oceans is quite low (Figure 7s) and they therefore do not contribute to the average over all transport directions much. On
 587 average, the updraught velocities increase from $\sim 0.4 \text{ m s}^{-1}$ up to $\sim 0.7 \text{ m s}^{-1}$ as the air masses approach SMEAR II. Slightly
 588 before arrival to SMEAR II (12–36 hours before arrival), difference can be observed in the updraught behaviour: winter
 589 updraught starts decreasing around 36 hours before arrival before increasing again at the 12-hour mark. During summer,
 590 the updraught increases all the way up ~ 18 hours, after which it steeply decreases and increases again at the same 12-hour
 591 mark as the wintertime updraught. As these differences are taking place relatively close to SMEAR II, it is likely that
 592 they are driven by the seasonal differences in the transport and local conditions very close to SMEAR II.

593 Activated fractions in ECHAM-SALSA display similar trends along their transport, increasing towards SMEAR II, but
 594 the seasonal difference in the magnitude is approximately 0.1, wintertime values being larger (Figure 7o). This difference
 595 stays nearly constant along the transport. Again, clear seasonal differences within the trajectory transport areas (Figure 7l
 596 and n) can be observed, and as the high activated fractions during winter (Figure 7n) do occur in high trajectory frequency
 597 areas (Figure 7s), they are more clearly reflected in the values when averaged over all transport directions (Figure 7o).
 598 As the seasonal differences N_{80} in ECHAM-SALSA are negligible, it is unlikely that the number of potential CCN is
 599 driving the seasonal differences in activated fractions and in the aerosol mass-precipitation relationships in Figure 4a.
 600 When the N_{act} is inspected (Figure S13t), however, somewhat larger number of particles have activated in winter
 601 compared to summer. Thus, when considering the large difference in the total number of particles (Figure S13e), the
 602 displayed differences in the activated fractions ($=N_{\text{act}}/N_{\text{tot}}$) are reasonable.

603 In addition to size, the chemical composition of the potential CCN also has an impact to their activation. The composition
 604 of Aitken and accumulation mode aerosols in ECHAM-SALSA (Figure S14b) does reveal, that the particles have
 605 relatively similar soluble accumulation mode SO_4 contribution in both seasons. The contribution of soluble OA in the
 606 accumulation mode is slightly larger in summer, but during winter, the smaller contribution from OA (in accumulation
 607 mode) seems to be compensated by larger contribution from soluble BC in the accumulation mode. Thus, the contribution
 608 from soluble modes altogether is relatively similar between the seasons and unlikely causes large differences in the
 609 particle hygroscopicity which could impact activation.

610 In order to investigate whether the seasonal differences in the activated fractions could also be due to slight differences
 611 in the sensitivity of activation to updraught velocities, we inspected the relationships between activated fractions and
 612 updraught velocities similar to UKESM1. For ECHAM-SALSA, the slope for summer is smaller (slope of 0.18, Figure
 613 8c) compared to winter (slope 0.36, Figure 8b). Thus, during winter, when the updraught increases, the activated fraction

can increase two times as much compared to summer. Therefore, despite the similar number of potential CCN in both seasons (N_{80} , Figure 7e), larger portion of those activate during winter, resulting to larger N_{act} (Figure S13t) and activated fractions (Figure 7o). All these findings discussed above are consistent with the stronger reduction for particle mass observed for ECHAM-SALSA in winter (compared to summer) in Figure 4a. During summer, very little to no reduction is observed for the particle number for ECHAM-SALSA in Figure 4b. The particle number concentration, however, is dominated by the small aerosols which are unlikely to activate (see also Figure S13e and Figure 2c). Therefore, even with high accumulated precipitation, no clear reduction is observed in Figure 4b during summer.

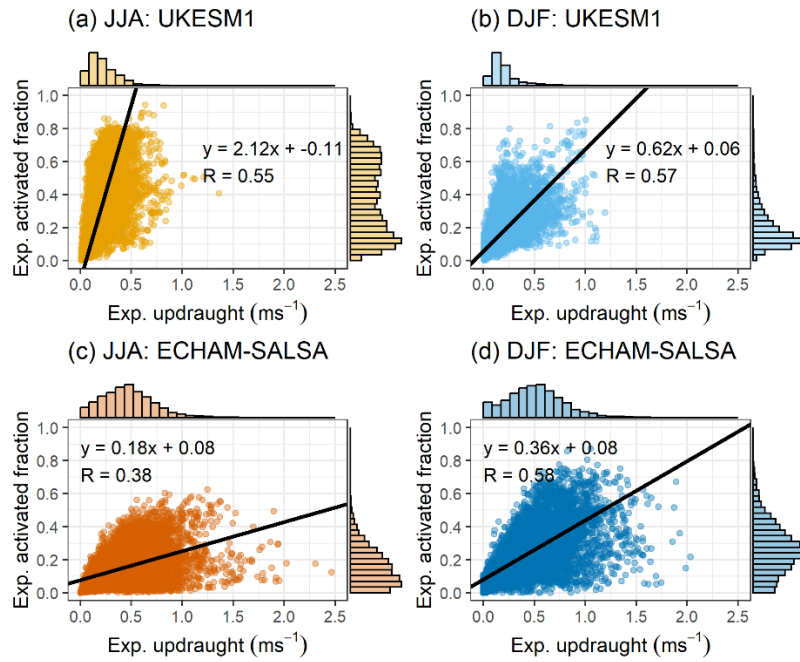


Figure 8 Average experienced activated fraction as a function of average experienced updraught velocity along the trajectories. Distribution of the values are shown with the histograms. JJA denotes summer (June-July-August) and DJF winter (December-January-February). Each coloured point denotes a median value determined from a single trajectory. The black lines show the regression line from orthogonal regression applied to the data shown and the legend show the slope, intercept and Pearson correlation (R) between the fit and the data. Note that the black regression lines extend over the whole plot area only due to visualization purposes.

4.4.2 Differences between GCMs and observations

Comparing the two GCMs in Figure 4 it is obvious that the seasonality in the aerosol-precipitation relationships is reversed: UKESM1 exhibits stronger reduction during summer but ECHAM-SALSA in winter. This is unlikely arising from the differences between the intensity of the precipitation during the travel of the air masses, as those are very similar between the GCMs (Figure S12a-e) within each season.

During summer, UKESM1 has less potential CCN (N_{80} , see Figure 7e) compared to ECHAM-SALSA, and also the updraught velocities are smaller in UKESM during summer, eventually leading to smaller number of cloud droplets too (N_{act} , Figure S13t). Comparison of the contribution of different chemical species in the accumulation (as these sizes have larger contribution to the particle mass) mode (Figure S14, top row), however, reveals that UKESM1 has much larger contribution of the soluble particles. This indicates, that during summer, the particles in UKESM1 have larger hygroscopicity, and could potentially activate more easily compared to ECHAM-SALSA. However, as the resulting N_{act}

(Figure S13t) in UKESM1 is smaller than in ECHAM-SALSA, the potentially larger hygroscopicity in UKESM1 particles do not seem to have significant impact on the droplet formation. When we consider the changes in the PNSD, however, where UKESM1 has significantly less particles but with larger average size compared to ECHAM-SALSA (which has more particles but smaller average size) as shown in Figure 2g and Figure S13e, it is sensible that larger activated fractions are observed for UKESM1 during summer as shown in Figure 7o. The difference in the activated fraction between the GCMs, however, is somewhat larger than what could be expected based on the differences in N_{tot} and N_{act} alone. Thus, also the relationships between updraught velocities and activated fractions were inspected to gain further insight. This reveals (Figure 8a and c), that indeed during summer, the slope between activated fractions and updraught velocities in UKESM1 is significantly larger (slope 2.12, Figure 8a) compared to ECHAM-SALSA (slope 0.18, Figure 8c)—difference being over 10-fold. This implies that even a small perturbation in updraught velocity in UKESM1 could increase the activated fraction drastically, resulting in the very high activated fractions observed in Figure 7o, despite UKESM1 having smaller updraught velocities in general. This could indicate a shift in UKESM1 cloud droplet formation from the updraught-limited regime to the transitional regime (e.g., Reutter et al., 2009). These findings align with the stronger reduction of particle mass in UKESM1 as shown in Figure 4a. The reduction of the observed particle mass in summer lies in-between of the two GCMs, initial reduction (up to 5 mm of accumulated precipitation) being more accurately represented by UKESM1.

The differences in the summertime reduction of particle number (Figure 4b) likely arise from the lack of boundary layer nucleation in UKESM1, thus affecting the number concentration of the smallest aerosol particles (see e.g., Figure 2g). As already discussed in Sect. 4.4.1, in SMEAR II, NPF is an important source of aerosols and the frequency of the NPF events has significant seasonal variation (Niemininen et al., 2014), summer and spring being most pronounced. Thus, the reduction of particle number in UKESM1 during summer (Figure 4b) is similar to the reduction of particle mass (Figure 4a), as both are dominated by relatively large aerosols. The summertime reduction of particle number in ECHAM-SALSA coincides with observations, which is to be expected as the Aitken and nucleation mode aerosol concentrations in ECHAM-SALSA are much closer to observed data than UKESM1 (Figure 2g and Table S5).

During winter, ECHAM-SALSA exhibits stronger reduction of particle mass compared to UKESM1 after ~5 mm of accumulated precipitation (Figure 4a). The N_{80} (Figure 7a-e) is relatively similar between the GCMs, but updraught velocities (Figure 7j) have large difference: UKESM1 updraught velocities range 0.2-0.4 m s⁻¹, whereas ECHAM-SALSA has values ranging approximately between 0.5-0.7 m s⁻¹. The higher updraught velocities in ECHAM-SALSA likely lead to the larger N_{act} (Figure S14t), thus eventually leading to the larger activated fractions for ECHAM-SALSA along most of the transport (Figure 7o) due to N_{tot} being relatively similar between the GCMs (Figure S13e) during winter. It should be noted, that the difference in activated fractions (Figure 7o) far away from SMEAR II is negligible. However, this difference drastically increases when air masses travel to SMEAR II: activated fraction in ECHAM-SALSA continues to increase while UKESM1 fractions stay nearly constant. Thus, it is unlikely that the similar activated fractions far away from SMEAR II significantly impact the reduction observed in Figure 4a.

Comparison of the particle chemistry in the accumulation mode in winter reveals that the GCMs have (Figure S14, bottom row) relatively similar fractions of soluble material. UKESM1 tends to have more SO₄, but ECHAM-SALSA more soluble OA and BC. In ECHAM-SALSA, however, the insoluble modes are not strictly insoluble but rather less insoluble compared to soluble modes (Sect. S2.3) and can thus also activate. This could lead to larger N_{act} (Figure S13o) and thus larger activated fraction (Figure 7o), considering that the difference in N_{tot} (Figure S13e) between the GCMs is clearly

smaller in winter than what it was in summer. The differences in the relationships between activated fractions and updraught velocities for the GCMs (Figure 8) are more subtle in winter (UKESM1 slope 0.62, ECHAM-SALSA slope 0.36) compared to the values in summertime discussed earlier. Activated fraction in UKESM1 does exhibit higher “sensitivity” for updraught velocities, however, due to the much larger updraught velocities in ECHAM-SALSA, this is likely not enough to increase the activated fraction to the same level, thus leading to less efficient reduction. These assessments align with the particle mass reductions in winter shown in Figure 4a, where particles at ECHAM-SALSA reach slightly lower end concentrations with high accumulated precipitation compared to UKESM1.

The differences in the wintertime reduction of particle number (Figure 4b) are less pronounced compared to those in particle mass (Figure 4a). Initial reduction seems to be more effective on UKESM1, however, after ~5 mm of accumulated precipitation, the reduction in ECHAM-SALSA becomes stronger. These differences between the GCMs, however, were not statistically significant (Kruskal-Wallis rank sum test, $p \geq 0.01$). The observational data exhibits stronger reduction than the GCMs during winter for the particle number (Figure 4b) up to ~10 mm of accumulated precipitation. After that, the observations overlap with ECHAM-SALSA. These inconsistencies could also arise from the fact that both GCMs have difficulties representing the bimodal particle number size distribution correctly during the winter months (Figure 2i).

4.4.3 Additional reasons for inter-model differences

Aside from differences driven by aerosol activation, it is important to note that during both summer and winter, additional factors can also contribute to the observed differences in the reductions (Figure 4). For example, the differences in the reduction of the particle mass (Figure 4b) could be influenced by the plume scavenging scheme, a feature only present in UKESM1 (see Sect. S2.4). In this process, aerosol activate into cloud droplets within the convective updraught and fall out via the main precipitation shaft of the cumulonimbus (Kipling et al., 2013; Mulcahy et al., 2020). Note that even though the particle mass is shown as a function of accumulated stratiform precipitation (Figure 4), the air mass trajectories have experienced convective precipitation too. Thus, removal via nucleation (which is more efficient for larger particles) followed by rainout in the convective plume, could also contribute. Inspection of the contribution of the precipitation types reveals that the contribution from the convective precipitation during summer is indeed slightly larger in UKESM1 compared to ECHAM-SALSA (Figure S15). This difference could be reflected in more effective summertime reduction in the particle mass in UKESM1. Another explanation for the more effective reduction of the aerosols during summertime in UKESM1 could be arising from the differences in the parametrizations of the re-evaporation of the falling droplets. In UKESM1, this process is not considered (see Sect. S2.3 and Mulcahy et al., 2020) whereas in ECHAM-SALSA evaporation of the droplets can occur and thus release the aerosols back to the atmosphere (e.g., Stier et al., 2005). During summertime, this re-evaporation could be enhanced due to higher temperatures, leading to less effective observed reduction of aerosols in ECHAM-SALSA compared to UKESM1. However, there can also be other explaining factors, such as location of the precipitation during travel, emissions and dry deposition, which could also indirectly cause differences between the models. Quantifying the exact processes from model parametrizations causing the differences between the observed relationships between aerosol mass and integral precipitation likely requires specific model sensitivity simulations to investigate this, thus being out of the scope of this study.

5 Lagrangian analysis on the effects of aqueous phase processing on aerosol chemical composition

In the analysis presented in this section, the relationship between the chemical processing occurring within clouds and fogs in the aqueous-phase is investigated. A special interest is in aqueous-phase SO_4 formation due to its high occurrence in the atmosphere (e.g., Ervens, 2015; Huang et al., 2019; Liu et al., 2020b). We employ a cloud proxy based on relative humidity (RH) along the trajectories similar to Isokääntä et al. (2022). To this end, the history of the air mass is investigated, and if the RH exceeds 94 %, we assume the air mass is in cloud. Further, the air masses were then separated into “clear sky” in which they had no experience of clouds or precipitation during the last 24 hours, and “in-cloud” when the RH exceeded 94 % at least at one trajectory point but no precipitation events occurred during the last 24 hours (Table S7). Only the last 24 hours of the air mass history were considered, as with longer air mass histories (i.e., longer investigated time) the number of strictly in-cloud trajectories decreases due to increasing possibility for precipitation events. Sensitivity tests were conducted by adjusting both the RH limit (from 90 % to 98 %) and trajectory length (from 12h to 60h), but they did not affect our conclusions. It was found that the trajectory length adjustment has large effect on the statistical reliability of the results, hence the investigation is limited to the last 24 hours and thus also stayed consistent with the previous investigation in Isokääntä et al. (2022). This approach is applied for ERA-Interim reanalysis and for the GCM trajectories in similar manner.

Reader should also note that UKESM1, ECHAM-SALSA and ERA-Interim do not necessarily have identical definitions for RH which could impact the results. To acknowledge this, we also investigated how well the RH along the trajectories actually describes the in-cloud cases by comparing this RH-based proxy to the co-located cloud fraction data from GCMs. This analysis is presented in Sect. S6, and overall, the cloud events (number of the events and their locations at the trajectories) from both approaches were similar, leading to similar conclusions as presented in Sect. 5.1 and 5.2 below. The precipitation used in the classifications here is the total precipitation (including both stratiform and convective precipitation), as aqueous-phase processes are taking place no matter the cloud type. Relative humidity data is from the HYSPLIT output instead of using raw GCM/ERA-Interim outputs with manual co-location. This is because UKESM1 was extracted on pressure levels instead of model levels, and the latter were used in this work for the manual co-location allowing consistency between other variables. The seasonal division applied here is based on the temperature, as in Sect. 4.2. To see whether transport directions and consequently the precursor emissions matter, data is divided into more clean and more polluted air masses (trajectories visiting latitudes below 60° north assigned to polluted sector as in Isokääntä et al., 2022). Trajectory frequency maps for these sectors are shown in Figure S16.

In this section, the variation in the total submicron mass of different chemical species depending on the experienced conditions is first examined and discussed for the GCMs (Sect. 5.1) and reflected to observations. Then, in the next section (Sect. 5.2), a size-resolved analysis is conducted to determine whether additional insight into in-cloud processing in GCMs could be provided.

5.1 Effects of in-cloud processing for total submicron aerosol mass

Both observations and GCMs show higher SO_4 mass concentrations for cloud-processed air masses within the “cold and polluted” (CP) sector (Figure 9), consistent with findings from Isokääntä et al. (2022). This pattern holds despite the reduced observational dataset due to temporal harmonization with the GCMs (see Sect. 2.4). Other air mass sectors are shown in the supplementary material (Figure S18).

751 Across all air mass sectors, both GCMs agree well with observations, considering expected differences in the total mass
752 concentrations. Statistically significant increases in SO₄ mass for in-cloud versus clear-sky air masses were found in both
753 observations and models ($p \leq 0.001$, Kruskal-Wallis test; Table S8), except for the warm and clean sector (Figure S17g-
754 f), where no clear difference was observed. As in Isokääntä et al. (2022), this may reflect limited SO₂ availability for
755 aqueous-phase oxidation in cleaner, warmer air masses. Supporting this, UKESM1 shows the lowest SO₂ levels in clean
756 sectors (CC and WC; Figure S18e), while higher SO₂ in polluted sectors (CP and WP) coincide with greater SO₄
757 differences. Recent findings from the Holuhraun eruption (Jordan et al., 2023) also suggest aqueous-phase oxidation
758 dominates SO₂-to-SO₄ conversion in GCMs. While future increases volcanic activity (Chim et al., 2023), could enhance
759 SO₂ levels and boost in-cloud SO₄ production, ongoing emission controls may reduce anthropogenic SO₂, potentially
760 counteracting this effect and influencing aerosol size and composition.

761 The observations shown here do not exhibit statistically significant differences for OA between the clear sky and in-cloud
762 air masses in any of the sectors. The median mass of OA in ECHAM-SALSA is larger for the in-cloud air masses for the
763 cold and polluted sector (Figure 9c and Table S8), but no other sectors exhibit statistically significant differences.
764 However, this difference in the OA mass in the cold and polluted sector is unlikely due to formation of aqSOA, as the
765 simulations employed in this study here did not explicitly model the formation of SOA. UKESM1 displays larger
766 differences in the OA mass, in which most are also statistically different. However, the same applies as for ECHAM-
767 SALSA, i.e., the model simulations do not include the formation of SOA, and thus the differences must arise from other
768 affecting factors. Both GCMs employ CMIP6 emission datasets as noted in the model setup for AeroCom Phase III GCM
769 Trajectory Experiment, and thus the differences observed here unlikely arise from varying emissions. One should also
770 keep in mind that the representations of OA in the GCMs might differ, and especially their relationship with temperature,
771 relevant driver for SOA formation in general, has been shown to exhibit large structural uncertainties between the GCMs
772 (Blichner et al., 2024).

773 Isokääntä et al., (2022) did not observe significant aqueous-phase SOA (hereafter, aqSOA) formation from the
774 observations and this has also been noted previously (Graham et al., 2020) for similar boreal environment. Formation of
775 SOA from gaseous precursors dominates this boreal region (see e.g., Petäjä et al., 2022), and thus distinguishing aqSOA
776 from the total formed SOA with our methodology is challenging. For isoprene-dominated environments, the formation
777 of aqSOA is a significant source for total SOA burden (e.g., Lamkaddam et al., 2021). Also biomass burning emissions
778 have been identified as a potential source for aqSOA (Gilardoni et al., 2016; Wang et al., 2024).

779 It was reported earlier that the observations also suggested increase in the mass fraction of SO₄ when the air masses had
780 been exposed to in-cloud conditions long enough (Isokääntä et al., 2022). To investigate whether similar behaviour could
781 be observed for the GCMs, we calculated the total time spent under the influence of non-precipitation clouds from the
782 96h long trajectories. Figure 10 demonstrates slight increases in the mass fraction of SO₄ with increasing time spent in
783 non-precipitating clouds for both GCMs. This, however, is somewhat affected by the data size. If inspecting the GCM
784 data which is temporally harmonised to the observations (Figure 10a-b), the conclusion is not as obvious compared to the
785 case were inspecting all available GCM data (Figure 10c-d). This highlights the importance of long enough GCM
786 simulations needed in this type of Lagrangian analysis utilizing single particle air mass trajectories unless ensemble
787 trajectories are utilised.

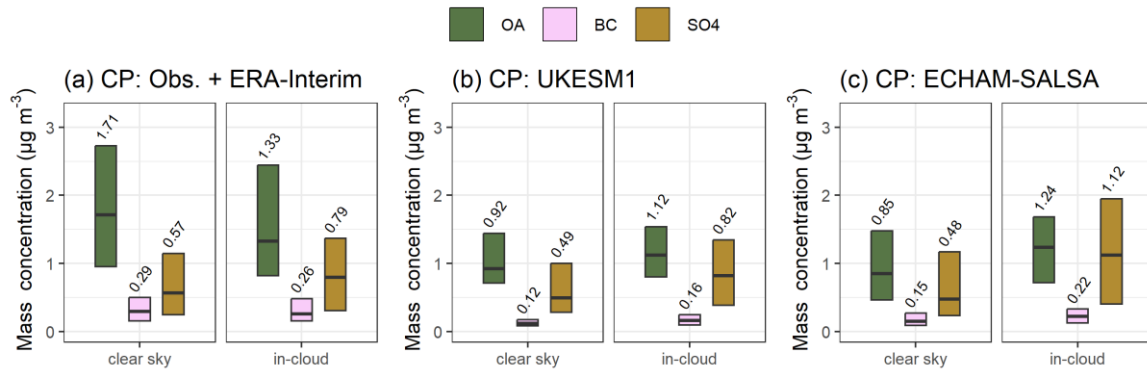


Figure 9 Median (black horizontal lines and numerical values) particle mass concentrations at SMEAR II with 25th–75th percentiles (boxes) for OA, eBC, and SO_4 for the cold and polluted (CP) air mass sector. The experienced conditions by the air mass are denoted as clear sky and in-cloud (non-precipitating). Subplots include (a) SMEAR II + ERA-Interim, (b) UKESM1 and (c) ECHAM-SALSA.

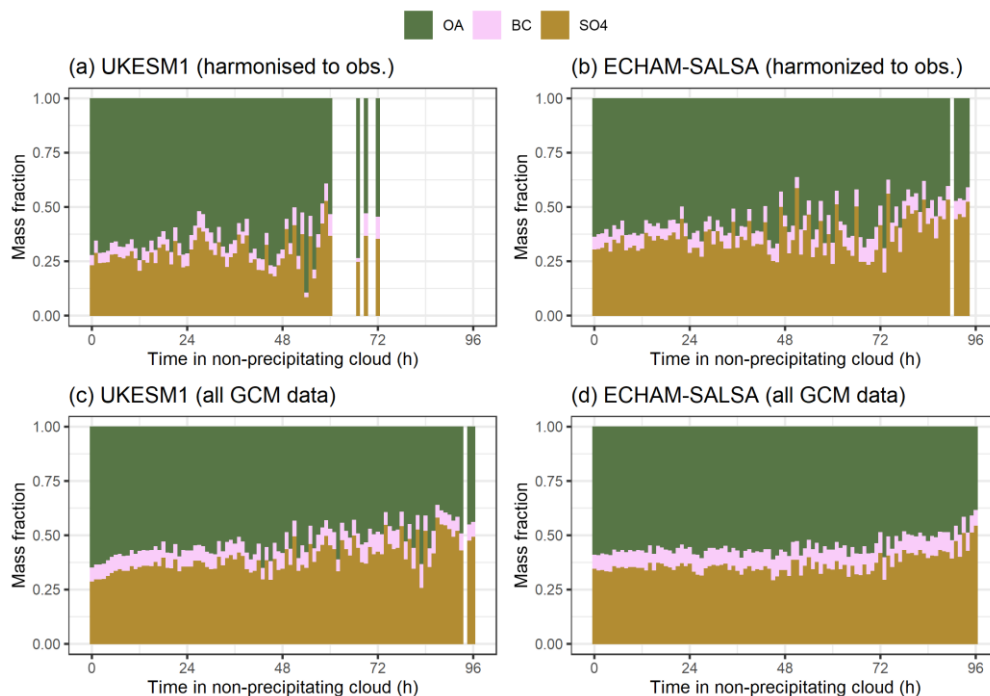


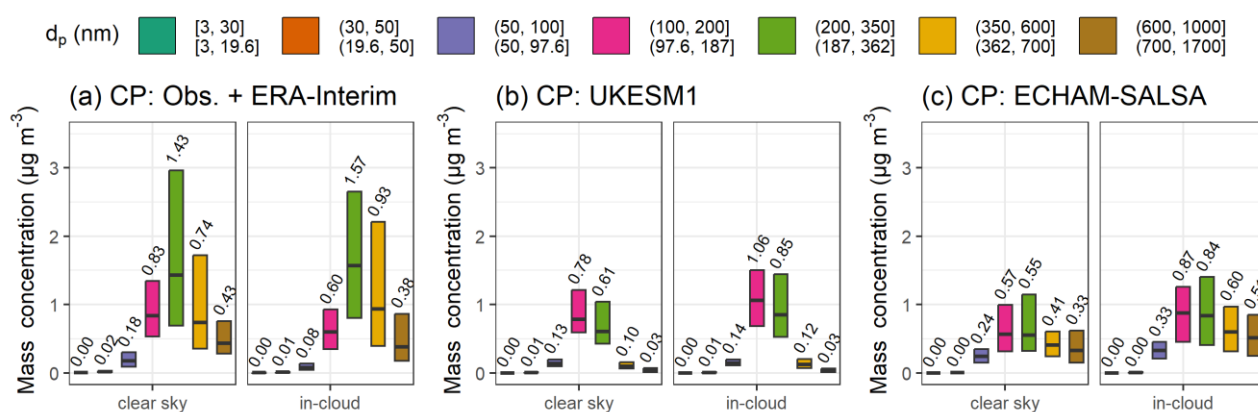
Figure 10 The mass fractions of OA, SO_4 , and BC for the more polluted air masses as a function of time spent in in non-precipitating cloud. The top row (a-b) shows the temporally harmonised data and bottom row displays the GCM data without harmonization. The figure shows mass fractions derived from median concentrations for each 1-hour bin.

5.2 Effects of in-cloud processing for size-resolved aerosol mass

To see whether the observed in-cloud formed SO_4 mass in the GCMs (Figure 9b-c) is contributing to same particle sizes as in the observations reported in Isokääntä et al. (2022), the analysis was repeated here for the GCMs. The observations indicated SO_4 mass originating from aqueous-phase processes is mostly contributing to particles with diameters of 200–1000 nm. Figure 11 shows the particle mass concentrations for various size classes derived from the PNSDs from the GCMs for the clear sky and cloud processed air masses for the cold and polluted sector. The three other sectors are shown in Figure S19, and Table S9 shows the results for the GCMs from the statistical significance testing between the clear sky

804 and in-cloud groups within each size class. Compared to observations, UKESM1 data (Figure 11a and Figure S19) implies
 805 the mass increase seems to be mostly distributed to bins with $d_p = 100$ -350 nm and up to 600 nm in the cold and polluted
 806 and cold and clean sectors. This is likely due to UKESM1 having large concentrations of particles in general within this
 807 size range (see e.g., Figure 2d). Like the observations, UKESM1 does not exhibit any mass increases for any of the size
 808 bins in the warm and clean sector (Figure S19e), being in line with no observed increase in the SO_4 mass in the same
 809 sector (WC) between the clear sky and cloud processed air masses (Figure S17h).

810 ECHAM-SALSA (Figure 11b and Figure S19), exhibits increased mass concentrations for sizes starting from $d_p = 50$ nm
 811 (only in cold and polluted sector) up to 1700 nm, depending on the sector. The largest bin here in ECHAM-SALSA might
 812 also be influenced by $d_p = 1$ -1.7 μm particles, which are neither considered in UKESM1 nor in the observations when
 813 inspecting the chemical components (see Sect. 2.4.2). Like UKESM1, ECHAM-SALSA also does not exhibit mass
 814 increases for any of the size bins for the warm and clean sector (Figure S21f).



815

816 **Figure 11 Median (black horizontal lines and numerical values) particle mass concentrations with 25th–75th percentiles (boxes)**
 817 **for selected size bins for (a) observations with ERA-Interim, (b) UKESM1 and (c) ECHAM-SALSA for the cold and polluted**
 818 **(CP sector). For the latter, the native size bins are shown (bottom row of the legend). The experienced conditions by the air**
 819 **mass are denoted as clear sky and in-cloud (non-precipitating).**

820 An advantage of the GCMs used in this study is their provision of size-resolved chemical composition, shown as mass
 821 fractions in Figure S20. For UKESM1, increase in the soluble SO_4 in the accumulation mode can be observed (Figure
 822 S20a). Due to the model structure, however, the accumulation mode itself consist of a large spread of particle sizes ($d_p =$
 823 100-1000 nm), i.e., internally mixed aerosols with external size modes, thus not providing additional information to our
 824 PNSD based analysis. For ECHAM-SALSA, the original sectional bins can be inspected (Figure S20c) thus
 825 corresponding to the PNSD bins presented in Figure 11b. All size bins that exhibited mass increases in Figure 11b also
 826 exhibit higher mass fraction for SO_4 in Figure S20c.

827 The observed changes in particle number size distributions (Figure 11) reflect the actual model parameterizations. In
 828 UKESM1, SO_4 produced via aqueous-phase chemistry is allocated to the soluble accumulation mode ($d_p > 100$ nm) and
 829 coarse mode ($d_p > 500$ nm) (Mann et al., 2010), with the results here showing increases in the 100–600 nm range. In
 830 ECHAM-SALSA, aqueous-phase SO_4 is distributed across soluble size bins spanning 50–10000 nm (2a bins; see Table
 831 S3, Bergman et al., 2012), with sector-dependent mass increases observed between 50–1700 nm. . In terms of aqueous-
 832 phase oxidation of SO_2 , both GCMs have similar parametrizations, and for example, oxidation of SO_2 by ozone (O_3) and
 833 hydrogen peroxide (H_2O_2) is considered in both (Bergman et al., 2012; Hardacre et al., 2021).

834 6 Conclusions

835 In this study we investigated the effects of stratiform precipitation (wet removal) and clouds (aqueous-phase oxidation)
836 on submicron aerosols along air mass trajectories. Two global climate models—UKESM1 and ECHAM-SALSA—were
837 analysed using a Lagrangian framework consistent with Isokääntä et al. (2022), now being seamlessly applicable to GCMs
838 (Kim et al., 2020). Our geographical focus was the SMEAR II station in Hyytiälä, Finland, and the surroundings,
839 representative of the boreal environment.

840 Our first objective was to investigate whether the trajectory-based relationships between aerosols mass, number and
841 precipitation vary between the observations and the GCMs. For aerosol mass, the derived removal for observations fell
842 between the two models across seasons. In contrast, aerosol number showed stronger divergence: while ECHAM-SALSA
843 and observations indicated minimal reduction, UKESM1 exhibited significant loss in particle number via precipitation,
844 likely due to a lack of small particles in the model. Supplementary analysis comparing a wider ensemble of GCMs
845 indicated that these two models were broadly representative, with their aerosol–precipitation relationships generally
846 falling near the middle of the inter-model spread. Reductions in organic aerosol (OA), sulfate (SO₄) and black carbon
847 (BC) also followed the same general trends as total aerosol, although with weaker seasonal differences. Our use of
848 normalized submicron mass and number as a function of accumulated precipitation proved effective in comparing removal
849 across models, though it lacks detail on particle size evolution—an important topic for future work.

850 Aerosol activation into cloud droplets followed by rainout appears to be the dominant removal process also in this study
851 being in line with earlier work. UKESM1 results further supported this showing nucleation followed by rainout as the
852 largest contributor. The seasonal differences in the observed removals within the GCMs were evaluated further by
853 inspecting key variables, such as, number of potential cloud condensation nuclei (N_{80}) and updraught velocities. The
854 seasonal differences we observed in these variables, along with changes in particle chemistry during the transport, were
855 consistent with the aerosol-precipitation relationships. The relationship between activated fractions and updraught
856 velocities shows opposite seasonal patterns in the GCMs: UKESM1 has a stronger summer correlation, while ECHAM-
857 SALSA's is stronger in winter, though its seasonal variation is smaller overall. This behaviour further explains the
858 differences between the aerosol-precipitation relationships in which ECHAM-SALSA showed similarity to observations.
859 We hypothesize that UKESM1's pattern may stem from the absence of boundary layer nucleation, resulting in fewer small
860 particles during summer. This emphasizes the need for better representation of particle number size distributions (PNSDs)
861 in GCMs.

862 Earlier studies (Isokääntä et al., 2022; Khadir et al., 2023) have noted that surface precipitation data, commonly used in
863 trajectory analyses, may not accurately reflect precipitation experienced by air masses at trajectory height. Here, we used
864 vertically resolved precipitation from UKESM1 and found that surface precipitation serves as a good proxy in this
865 environment, where trajectories largely remain within the mixed layer and stratiform precipitation dominates. However,
866 this analysis only considered liquid precipitation and may not apply to regions where convective precipitation is more
867 prevalent. In such environments, the vertical distribution, intensity, and frequency of precipitation could differ
868 substantially, potentially altering the accumulated wet removal along trajectories. Therefore, while our results are
869 representative of boreal regions with stratiform precipitation, further work is needed to assess how applicable they are to
870 regions with different precipitation regimes.

871 Our second objective was to investigate whether the GCMs exhibit similar increase in sulfate mass due to in-cloud
872 production as the observational data. Both GCMs exhibited statistically significant difference in the SO₄ mass when air

873 masses with only clear sky experience were compared to in-cloud processes air masses. The SO₄ mass was larger for the
874 cloud processed air masses for all other air mass sectors (based on temperature and direction) except the warm and clean
875 air masses, where GCMs showed no significant difference between clear sky and in-cloud air masses. These results were
876 consistent with earlier work (Isokääntä et al., 2022). Availability of the SO₂ to be oxidised is likely determining whether
877 we see in-cloud production of SO₄, and from UKESM1 this was further supported by the SO₂ concentrations and their
878 seasonality. The size-resolved analysis reflected the model parametrizations, the aqueous-phase SO₄ being mostly
879 distributed in the larger aerosol sizes.

880 As expected based on Isokääntä et al. (2022), we did not observe significant aqueous-phase SOA formation. This is likely
881 due to the studied environment (boreal forest), and has also been noted previously (Graham et al., 2020) for similar boreal
882 forest environment. However, some increases in OA mass were seen in the GCMs despite the fact that aqSOA formation
883 was not explicitly modeled, possibly reflecting other processes or model inconsistencies. A recent study from Blichner
884 et al. (2024) also pointed out the large differences between GCMs concerning their OA-temperature relationships, which
885 could also contribute to the discrepancies observed here.

886 Overall, both GCMs reproduced the observed exponential decrease in aerosol mass with increasing precipitation and
887 showed similar cloud-processing behaviour for SO₄. Yet key seasonal differences remain, especially in aerosol–
888 precipitation relationships and their underlying drivers. A primary model bias identified in this study is the difference in
889 aerosol number size distributions compared to observations, particularly the underrepresentation of small particles in
890 UKESM1. Our results suggest that discrepancies arise more from differences in aerosol size distributions and updraught
891 velocities than from the wet removal parametrizations themselves. These variables also affect activated fractions and
892 cloud interactions, and they are shaped by processes beyond the 4-day analysis window.

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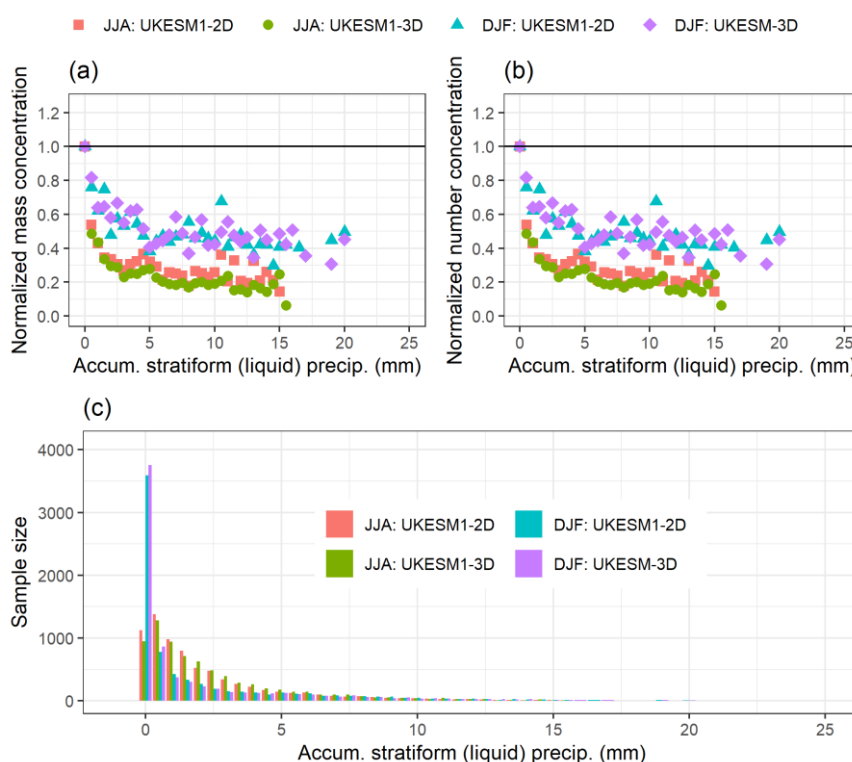
894 **7 Outlook**

895 While our results show encouraging agreement between observations and GCMs in overall aerosol–precipitation
896 relationships, key differences—especially related to seasonality and aerosol number—highlight the need for further work.
897 Future studies should investigate the evolution of aerosol size distributions along air mass trajectories in more detail and
898 better disentangle gas-phase and aqueous-phase sulfate formation. Expanding analyses to regions with dominant
899 convective precipitation is also important, as the findings here are limited to stratiform, liquid-phase conditions typical of
900 boreal environments. Including a wider range of GCMs, despite the computational demands, would help clarify the
901 structural causes behind the differences observed. Together, these efforts are essential for improving the representation
902 of aerosol–cloud–precipitation interactions in climate models.

903 Appendix A

904 The lack of vertical resolution in the precipitation data from ERA-Interim reanalysis or Global Data Assimilation System
 905 (GDAS, (<http://ready.arl.noaa.gov/archives.php>, last access: 3.2.2024) in studies using Lagrangian approaches is now
 906 being recognised (Dadashazar et al., 2021; Isokääntä et al., 2022; Khadir et al., 2023). Unfortunately, vertically resolved
 907 precipitation data from reanalysis datasets or GCMs, with high enough time resolution to be useful for trajectory models,
 908 is not a commonly provided diagnostic. For UKESM1, this diagnostic can be extracted. Here, we conducted a comparison
 909 between the vertically resolved and surface precipitation data along the air mass trajectories to investigate how well the
 910 surface precipitation describes the actual experienced precipitation by the air mass. Only liquid (stratiform) precipitation
 911 is inspected, as vertically resolved snowfall was not included in the variable extraction with high enough vertical
 912 resolution for this model run.

913 We first inspected the relationship between the normalized particle mass and number with the accumulated stratiform
 914 precipitation, similar to Figure 4. This assessed whether aerosol–precipitation relationships differ between surface and
 915 vertically resolved precipitation. Displayed in Figure A1, the results indicate the effects of stratiform precipitation at the
 916 height of the air mass are similar to the effects of stratiform precipitation at the surface. This is likely related to the average
 917 altitude of the air masses, as for SMEAR II they tend to travel well below the top of boundary layer.

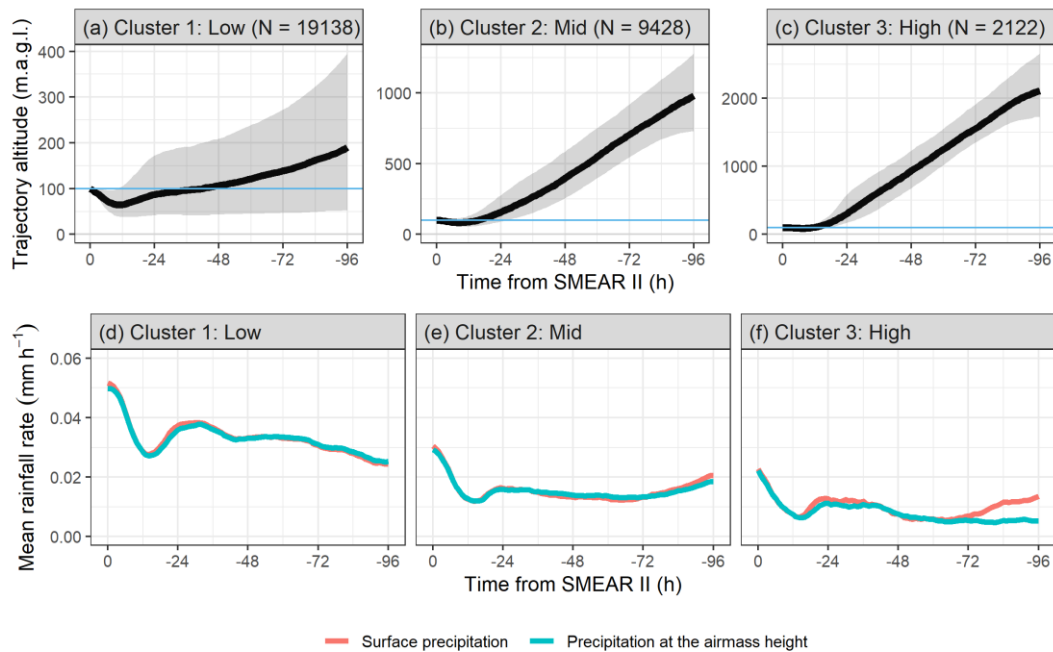


918 **Figure A1 Normalized total ($d_p = 3-1000$ nm) particle mass (a) and number (b) at SMEAR II for summer (JJA) and wintertime (DJF) as a function of 0-25 mm of accumulated liquid stratiform precipitation along the 96-hour long air mass trajectories at the height of the air mass (referred as 3D) and at the surface (referred as 2D) for UKESM1. The coloured points show the median values for each 0.5 mm bin of accumulated precipitation when the number of trajectories in the bin was 10 or larger. The sample size for the corresponding bins is shown in (c).**

924 To investigate whether the height of the air mass plays a role, as speculated in Isokääntä et al. (2022), the air mass
 925 trajectory altitudes were clustered with Kmeans (e.g., Hartigan and Wong, 1979) and 3 clusters with distinct height

926 profiles were selected for further analysis. Clustering each season separately provided similar height profiles as clustering
 927 of the whole data, and thus the latter approach is presented.

928 Figure A2 shows the median altitudes of the clusters and the corresponding mean stratiform rainfall rates. Overall, the
 929 mean rainfall rates show similar values despite the precipitation diagnostic. In the low-altitude cluster (Figure A2d),
 930 overall highest rainfall rates (mean over all trajectories and hours for surface precipitation, $\sim 0.033 \text{ mm h}^{-1}$) are observed.
 931 In the mid-altitude cluster, rainfall rates are smaller ($\sim 0.016 \text{ mm h}^{-1}$) compared to the low-altitude cluster, and in the
 932 high-altitude cluster, the rainfall rates are the smallest ($\sim 0.010 \text{ mm h}^{-1}$). In the high-altitude cluster (Figure A2f) more
 933 differences emerge between the two precipitation types, especially afar from SMEAR II.



934

935 **Figure A2 Clusters based on air mass trajectory altitudes for UKESM1. In (a)-(c) the black lines show median trajectory**
 936 **altitude as a function of time from SMEAR II and 25th to 75th percentiles are shown with the shaded area. The used arrival**
 937 **height at SMEAR II given to HYSPLIT is indicated with blue horizontal line. The corresponding mean rainfall rates are shown**
 938 **in (d)-(f). Clusters are named based on the maximum altitude the trajectory has resided during the last 4 days. Note the different**
 939 **y-axis limits in subplots (a)-(c).**

940 Each cluster was then further separated by season. The median altitudes, if inspected separately for each season, are nearly
 941 identical between the seasons within each cluster, and thus not shown here. Figure S21 shows the differences between
 942 the mean liquid rainfall rates between surface and vertically resolves stratiform precipitation (positive difference
 943 indicating the rainfall rates at the surface are higher) for each cluster and each season.

944 During autumn (SON) the two approaches for the precipitation exhibit observable differences only in the high-altitude
 945 cluster, where the surface precipitation shows some overestimation of the actual experienced precipitation by the air mass
 946 with increasing trend when moving farther away from SMEAR II. This could imply that the air mass has spent some time
 947 above or inside the precipitating cloud, as also the air mass altitude increases when moving away from the station (Figure
 948 A2a-c). During summer (JJA), all clusters mostly show precipitation at the air mass height being larger than the surface
 949 precipitation, expect in the high-altitude cluster (Figure S21c) 72 to 96 hours before arrival to SMEAR II. As the
 950 temperatures during summer are higher than in other seasons, this could be indication of evaporation as the surface

951 precipitation in UKESM1 includes only precipitation that reaches the surface i.e., it is not column integrated. During
952 spring (MAM) and winter (DJF) the surface precipitation shows small overestimation at some points along the trajectories,
953 and the differences are largest at the high-altitude cluster—where, however, the rainfall rates are very small overall (see
954 Figure A2f) for both precipitation types.

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958 **Data availability**

959 Raw observational data were collected by INAR, University of Helsinki. Field data (particle number size distributions
960 and black carbon) are freely available from <https://smear.avaa.csc.fi/download> (last access: 20 February 2022; Ministry
961 of Education and Culture of Finland and CSC, 2022). The ACSM data on aerosol composition are freely available from
962 the EBAS database at <http://ebas.nilu.no/> (last access: 20 February 2022; NILU, 2022).

963 The pre-processed observational data, ERA-Interim and GCMs trajectories along with the co-located variables used in
964 this study can be found from Talvinen et al., (2025b).

965 **Code availability**

966 Data analysis was conducted in R statistical software (R version 4.2.0, R Core Team, 2019) and Python (version 3.10.4),
967 and colour maps for the figures considering colour vision deficiencies were inspired by Crameri et al., (2020).

968 The scripts to reproduce the main findings both in R and python can be found from Talvinen et al., (2025a). Python scripts
969 for the data conversion (GCM output into ARL) and co-location of the GCM and reanalysis data variables to the
970 trajectories can be obtained from DGP.

971 **Author contribution**

972 DGP and AV proposed the study. ST, DGP and AV designed the research questions. ST had the lead role in data analysis
973 with supporting contribution from PK, DGP, ET and RC. The modelling framework to calculate trajectories from GCM
974 meteorological fields was conceived and performed by DGP with support from ZK and JT. The development and
975 application of this framework to the AeroCom GCMTraj model submissions was performed by PK with support from
976 DGP. Model simulations and data submissions were performed by ET, DGP, TK, EH, HK, TK, DN, DWP, YY, JZ and
977 SK. UKESM1 model simulation configuration was supported by AS, and ZK supported the processing of ERA-Interim
978 reanalysis data. HYSPLIT trajectories were calculated by PK and the co-location scripts were developed by PK with
979 supporting contribution from DGP, ET, ST and RC. Co-location of GCM data and ERA-Interim precipitation to the
980 trajectories were performed by ST. Results were interpreted by ST, DP and AV with supporting contribution from all co-
981 authors. The manuscript was written by ST with supporting contribution from DP. All co-authors commented, edited and
982 gave feedback on the manuscript.

983 **Competing interests**

984 At least one of the (co-)authors is a member of the editorial board of Atmospheric Chemistry and Physics.

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