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Ideas and Perspectives: Max MACS – constraining the potential global scale of Marine Anoxic Carbon Storage for CO_2 removal

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Abstract. Marine Anoxic Carbon Storage (MACS) is a potential strategy for enhancing atmospheric CO_2 removal (CDR) by sequestering organic carbon produced by terrestrial plants in stable, anoxic marine reservoirs. Initial results suggest that MACS could, in theory, operate at the gigatonne scale that would be required to impact global climate, with limited environmental risk and promising opportunities for co-benefits. However, several outstanding knowledge gaps make it challenging to quantify the actual potential global scale of MACS with confidence. To inform decisions about climate mitigation and trade-offs in the future, it is essential that we know how MACS implementation at scale would impact critical environmental and economic systems in the context of likely future scenarios.

Building on the results of a workshop in Bucharest, Romania in 2025, we discuss the potential impacts of MACS activities on the ecology, biogeochemistry, economy, and community around the Black Sea, seafloor brines, and other anoxic marine sites. Quantifiable limits to the potential maximum feasible scale of MACS for CDR are organized into five criteria: (1) Durable storage site capacity; (2) Biomass sources and logistics; (3) Greenhouse gas balance; (4) Oxygen and sulfide impacts at the redoxcline; and (5) Impacts on dissolved organic matter or nutrients in the oxic zone. For each criterion, we evaluate the factors that could limit scale, our current state of knowledge, and the priority knowledge gaps that, if addressed, would improve our ability to estimate the potential global scale of MACS for CDR. Research is needed to understand its potential impacts at scale, but MACS is nonetheless worthy of serious consideration as a potential pathway for climate mitigation in coming decades.

1 Motivation for Marine Anoxic Carbon Storage (MACS)

Marine Anoxic Carbon Storage (MACS) is a carbon dioxide removal (CDR) approach that would sequester organic carbon in the form of plant biomass in naturally-occurring, anoxic (O_2-free) marine environments like isolated basins, brines, or sediments. MACS is thus a hybrid land–ocean approach for CDR. Terrestrial, lignocellulosic organic matter can be efficiently preserved in these environments by natural mechanisms over timescales of thousands of years or more, preventing the reoxidation of this organic C to CO_2. Here, we focus exclusively on the potential sequestration of terrestrial plants; the use of aquatic biomass would introduce additional categories of potential impact on the ocean that are beyond the current scope (c.f., Roberts et al., 2025).

1.1 Earth's carbon fluxes and global CO_2 removal goals

Climate change is driven largely by excess atmospheric CO_2 (IPCC AR6, 2021). Given current emissions trajectories, on the order of 10 Gt per year of CO_2 (200 Gt CO_2 by 2050) will need to be removed from the atmosphere in coming decades to meet climate targets of ≤2°C average warming by the end of the century (IPCC AR6, 2021). Net greenhouse gas removals

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

can be met through a combination of approaches, each of which might contribute the equivalent of ~0.1 to 1 Gt worth of CO_2 removal (CO_2e) per year (National Academies of Sciences, Engineering, and Medicine, 2022). This scale of activity would approach that of some of the largest industries on the planet today, including cement and oil production, as well as that of modern fluxes in the global carbon cycle (Table 1). For example, global rates of organic carbon (OC) burial in marine sediments today are equivalent to 2.87 ± 1.3 Gt CO_2e/yr, including approximately 0.23 – 0.33 Gt CO_2e/yr of terrestrially-sourced OC (Dunne et al., 2007; Talling et al., 2024).

Process	Scale of Annual Flux (Gt/yr)			Source
	total mass	carbon	CO_2e	
Human Activities				
Global cement production (portland + masonry)	4.1			USGS, 2024
Global municipal solid waste processing	2.1			UNEP, 2024
Global oil production (crude oil + lease condensate)	4.48			US EIA, 2023
IPCC target for 2050+ (all CDR methods)		2.7	10	IPCC AR6, 2021
Moderate / high scalability target (each pathway)		0.03 / 0.3	0.1 / 1.0	US NASEM, 2022
Conventional CDR via afforestation		0.52 (0.33 – 0.6)	1.9 (1.2 – 2.2)	Vaughan et al., 2024
Modern Global Fluxes				
Global marine NPP		51.9 ± 3.3	190.3 ± 12.1	Zhang et al., 2025
Global terrestrial NPP		55.8 ± 2.1	204.6 ± 7.7	Zhang et al., 2025
NPP in the Black Sea		0.05 – 0.07	0.18 – 0.26	Demidov, 2008
Change in terrestrial OC stocks 1993–2019		1.3 ± 0.5	4.8 ± 1.8	Bar-On et al., 2025
River flux of particulate OC to the ocean		0.17 – 0.2	0.63 – 0.73	Burdige, 2005
OC burial in marine sediments		0.78 ± 0.36	2.87 ± 1.30	Dunne et al., 2007
Terrestrial OC burial in marine sediments		0.062 – 0.09	0.23 – 0.33	Talling et al., 2024
Earth History Analogs				
Excess OC burial in the ocean, OAE-2		~0.1	~0.3	Owens et al., 2018
Excess OC burial in the ocean, PETM recovery		0.083	0.30	Papadomanolaki et al., 2022

Table 1: Proposed scale of global CDR relative to natural and anthropogenic fluxes. NPP = net primary production; OC = organic carbon; OAE-2 = Ocean Anoxic Event 2; PETM = Paleocene–Eocene Thermal Maximum. Sources: USGS = United States Geological Survey; UNEP = United Nations Environment Programme; US EIA = United States Energy Information Administration.

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Terrestrial plants are active participants in the inter-annual carbon cycle and their growth has substantial leverage over atmospheric CO_2 concentrations. Annual fluctuations in the Keeling Curve show seasonal oscillations in atmospheric CO_2 of ~5 ppm, which is roughly equivalent to ~ 39 Gt CO_2e (Ballantyne et al., 2012). A slight imbalance between terrestrial production and respiration allows nearly a third of current anthropogenic C emissions to be stored on land (Friedlingstein et al., 2025). Recent results suggest that most of this storage may be in the form of nonliving (soil) organic carbon; models suggest that global stocks may have increased by as much as 4.8 Gt CO_2e/yr since 1993 (Bar-On et al., 2025). For comparison, conventional CDR (e.g., afforestation) currently uses terrestrial plants to remove ~2 Gt CO_2e/yr (Vaughan et al., 2024).

Throughout Earth's history, changes in the fluxes of organic carbon into long-lived reservoirs like soils and marine sediments have substantially impacted atmospheric CO_2 concentrations and global climate, typically over timescales of thousands to tens of thousands of years. For example, during events like Ocean Anoxic Event 2 (OAE-2; ~97 Ma), widespread anoxia supported the burial and preservation of vast amounts of largely marine OC, which contributed to global cooling (Jarvis et al., 2011). Marine sediments accumulated approximately 40,000 Gt excess C (150,000 Gt CO_2e) over the course of this 500-kyr event (Owens et al., 2018), which implies an average excess OC burial of ~0.3 Gt CO_2e/yr. This average presumably includes shorter periods of higher excess OC burial, but it is nevertheless of the same order of magnitude as scalability goals for individual CDR pathways (NASEM 2022, Table 1).

Terrestrial OC burial has also contributed to more recent shifts in global climate. Fluxes of terrestrial OC to the ocean via rivers are sensitive to changes in hydrology and weathering, both of which generally increase under warmer climates, generating a stabilizing climate feedback. Under the right conditions, the preservation efficiency of terrestrial materials in coastal sediments can approach 100%, although it averages 20–44% today (Galy et al., 2007; Blair and Aller, 2012). During the Paleocene–Eocene Thermal Maximum (PETM) at ~56 Mya, the burial of terrestrial OC in marine sediments may have contributed substantially to enhanced global carbon burial and CO_2 drawdown (Inglis et al., 2025). Approximately 3,300 Gt excess OC was buried in marine sediments during the 40,000-year recovery period, equivalent to an average sequestration rate of around 0.30 Gt CO_2e/yr (Table 1; Papadomanolaki et al., 2022). Terrestrial OC burial in marine sediments may have contributed similarly to recovery from Eocene hyperthermals and other hothouse events with high rates of weathering (Inglis et al., 2025). Average rates of excess global OC preservation associated with periods of climate cooling in the Cenozoic thus approach the scales of sequestration being proposed for individual CDR pathways (Table 1). Like OAE-2, these events can serve as valuable analogs to understand the mechanisms and impacts of enhanced OC preservation in the ocean.

1.2 Current state of MACS research and development

A variety of different CDR approaches will be needed to meet total CDR targets (NASEM, 2022), but potential techniques vary substantially in terms of their scalability, energy efficiency, carbon storage durability, economic implications,

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

and environmental and social impacts. Different CDR approaches are expected to come "on-line" at different times in the coming century (Ganti et al., 2024), ideally in a way that would optimize for their relative advantages under different scenarios.

MACS has several potential advantages that make it a promising target for further research. Because it relies on plants to concentrate atmospheric CO_2, MACS avoids the need for "clean" grid power that plagues direct air capture and other energy-intensive CDR approaches. Additionally, biomass sourcing can rely on largely extant agricultural infrastructure, minimizing the CO_2 emissions from construction. MACS could therefore have some of the lowest energy requirements per tonne and some of the highest carbon efficiencies of any proposed CDR approach (Strand and Benford, 2009; Sanchez et al., 2025). Marine anoxic storage is only one of several proposed methods for sequestering terrestrial biomass; woody materials are also being stored in anoxic vaults on land (Zeng et al., 2024) and as slurries in deep wells (Snyder, 2022). Carbon efficiency and energy needs for any of these storage alternatives will depend on biomass type, transportation mode, and processing method, which is likely to favor different storage options in different locations. The first carbon credits for terrestrial woody biomass burial on land were issued in 2024 (https://registry.puro.earth/issuances).

Although biomass materials described here are sourced from land, their long-term storage would occur in specific parts of the ocean. MACS differs from other ocean-based methods for C storage like ocean alkalinity enhancement (OAE), direct ocean carbon capture and storage (DOCCS), or marine biomass cultivation in two notable ways: first, it sources plants, which directly remove carbon from the atmosphere without requiring air-sea gas exchange; and second, relatively isolated reservoirs are targeted for storage. Spatially limited sequestration sites can facilitate monitoring and permitting within a single jurisdiction, and they may reduce risks of impact to surrounding ecosystems. Nevertheless, MACS faces many of the same challenges as other marine CDR techniques, especially related to legal, social, and political issues.

The intentional placement of materials in the ocean, whether that is termed storage, burial, or dumping (e.g., in the U.S. EPA MPRSA framework), is regulated by the London Convention and Protocol. Active negotiations are underway to update the London Protocol framework to be fit-for-purpose for ocean-based CDR, and this legal framework is likely to continue to evolve rapidly in coming years (GESAMP, 2025; Silverman-Roati and Webb, 2025). Within current legal frameworks, early field studies are underway, but regulatory barriers and permitting challenges remain significant. In the United States, the Environmental Protection Agency (EPA) announced in June 2025 that they had "tentatively determined" that they would issue a research permit that would involve the placement of 20 tonnes of sugarcane bagasse into Orca Basin, a hypersaline brine pool on the deep (2,200 m) U.S. continental shelf, over 18 months (epa.gov/carboniferous). In Europe, the Romanian National Environmental Protection Agency (ANPM) requested the execution of an Environmental Impact Assessment prior to permitting the placement of 100 tonnes of woody biomass on the deep Black Sea seafloor.

In addition to establishing clear and efficient regulatory pathways, successful growth of the MACS field will require mechanisms for verifying the results of CDR actions in a way that earns social acceptance as well as trust in the scientific community. One step toward that goal is the release of a first protocol for future carbon credit verification this year (Puro.Earth,

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

2025). This protocol focuses exclusively on enclosed anoxic basins and lignocellulosic (woody) biomass, and it requires both ex-situ experiments and in-situ measurements as evidence for ≥200 yr storage. Over the long term, revised protocols could ideally be embedded into regulatory frameworks and environmental impact standards such as the Marine Strategy Framework Directive or the EU Carbon Removal Certification Framework.

1.3 Scope and Purpose of this Analysis

MACS research currently sits at an inflection point. There is a strong infusion of energy into the CDR field as awareness grows and initial scientific results are published in the peer-reviewed literature. At a time of extraordinarily tight resources for climate-related research and development, however, the justification for serious investment in MACS as a CDR pathway hinges on the question of whether this process can actually make a serious dent in atmospheric CO_2. Currently, we don't know the answer – but we do know how to get there.

The purpose of this paper is to review the current state of knowledge related to predicting the outcomes of MACS at a climatically-relevant scale and to identify gaps in that knowledge. We aim to integrate global perspectives as a step toward building international research capacity and collaborative networks for MACS-related science, and to support and develop an international research infrastructure and network capable of testing feasibility on a global scale.

This work was catalyzed by a two-day intensive workshop in Bucharest, Romania in February 2025, which involved primarily academic and NGO researchers but also experts in governance, CDR verifiers, local environmental leaders, and entrepreneurs. Participants in the workshop came from 14 countries and 25 institutions, including all countries with Exclusive Economic Zones (EEZs) relevant to the Black Sea and Orca Basin except Russia and Ukraine. The goals of the workshop were to broadly discuss the CDR potential and risks of MACS with local and global scientists and stakeholders, and to identify high-impact research questions that could help evaluate limits to scale. The central ideas in this manuscript arose from workshop discussions, as refined through iterative conversations over the subsequent months.

Philosophically, the authors emphasize that support for research is not an endorsement of any particular technology or action at scale. Participants in the Bucharest Workshop came to the consensus conclusion that there may be significant potential for MACS to contribute to global CDR needs, and that the risks associated with small-scale field research are expected to be negligible relative to its potential value. They also concluded that several clear uncertainties limit our ability to predict some important impacts of MACS implementation at scale. Addressing these uncertainties would provide critical information for future decision makers in government, philanthropy, and industry. This report represents the first part of a roadmap to answer those questions.

The authors also acknowledge that the maximum potential global scale of MACS is only one component of a societal strategy toward responsible decision making around climate mitigation and CDR. This analysis is intended to support

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

interdisciplinary efforts by identifying quantifiable physical and biogeochemical limits on the potential scale of MACS, which can help guide resources toward strategies with the potential for meaningful climate impact and inform ongoing conversations around how to meet global needs for hundreds of gigatonnes of net CO_2 removal in coming decades.

2. Defining potential locations for MACS

In order to calculate a realistic estimate for the maximum feasible scale of MACS for CDR, we need to first decide which anoxic environments to prioritize in our analysis.

Anoxic marine environments in the water column and sediments are extremely diverse, including deep-sea brines, inland seas, river deltas, and large regions of sediments along the continental shelves. Participants at the Bucharest Workshop discussed the benefits and downsides of including different types of sites in a MACS global scale estimate. The Black Sea and deep hypersaline anoxic basins (e.g., Orca Basin) were selected for initial analysis due to their long-term redox stability and intense stratification, as well as their relatively mature level of scientific understanding and operational readiness. Carbon storage in a third type of anoxic environment, rapidly-accumulating sediment at (e.g.) major river mouths, is less well understood but has the advantages of engaging multiple sites across the global North and South and having vast potential storage capacity. Although this report prioritizes the Black Sea and brines, it retains a consideration of sediment burial as a forward-looking area for potential expansion.

2.1 The Black Sea

The Black Sea is the world's largest permanently anoxic basin, with a sulfidic water column below its shallow chemocline at approximately 80 to 150 meters depth. Abyssal waters between 1,750 – 2,212 m have a volume of ~75,137 km^3, most of which have been isolated from the atmosphere for >1,000 years (Murray et al., 1991; Lee et al., 2002). Conditions in the Black Sea have famously preserved wooden ships for millennia (Pacheco-Ruiz et al., 2019). Its isolation and long water residence time make the Black Sea a critical site for MACS at scale, particularly as it exists in a region with abundant potential biomass sources and transport infrastructure. However, there is also a long history of conflict and environmental damage in the region, such as ongoing warfare and the dumping of nuclear materials.

For MACS, the leading environmental concerns with this site pertain to its shallow redoxcline that separates oxic from anoxic waters. This interface and the overlying oxic waters are critical to communities in the region for fisheries, recreation, and cultural and historical significance. The position of the redoxcline has historically moved due to changes in water and carbon budgets, regional climate, and eutrophication (nutrient input), requiring coupled models to project its behavior into the future (Ilicak et al., in prep). These factors also impact ecosystems throughout the upper oxic zone of the Black Sea (Mikaelyan et al., 2013) and control whether surface waters are a net source or sink for CO_2 (Sergeev et al., 2024).

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

Establishing the current sensitivity of the Black Sea system to changes in drivers like surface temperature, nutrient inputs, river flows, and circulation will be essential to understand its baseline state and facilitate comparisons against the outcomes of potential MACS interventions.

2.2 Orca Basin and other brines

The Orca Basin is a 10.24-km^3 brine pool located at 2,200 m depth in the Gulf of Mexico, approximately 200 km from the Louisiana coast (Shokes et al., 1977). It formed roughly 7,800 years ago when a buried layer of salt was exposed on the seafloor and dissolved. The brine is strongly inhibited from mixing with the overlying seawater due to its strong density gradient (Addy and Behrens, 1980). As a result, Orca Basin is anoxic and physically traps both solid and dissolved carbon (Harvey and Kennicutt, 1992; Shah et al., 2013)). Sediments from Orca Basin exhibit exceptional organic matter preservation efficiencies, with reports of unprecedented levels of seaweed preservation for thousands of years (Kennett and Penrose, 1978). Other brines are known across the U.S. continental shelf, the Mediterranean Sea, and the Red Sea, although only a few examples have received as much research attention as the Orca Basin.

2.3 River deltas and research sites

Major river deltas like those of the Amazon, Mississippi, Ganges-Brahmaputra, Yangtze, and Congo transport hundreds of millions of tonnes of sediment each year into coastal and offshore environments. River fans today bury the equivalent of 0.18–0.37 Pg/yr CO_2e, and they have been important carbon sinks throughout Earth's history (Blair and Aller, 2012). Due to the combination of rapid sedimentation rates and high organic matter concentrations, river fan sediments often develop anoxic zones that contribute to total organic carbon preservation.

In the Bay of Bengal region, for example, monsoonal rivers are a primary driver for land-to-sea transport of plant- and soil-derived organic carbon (Contreras-Rosales et al., 2016). Water and sediments are carried into the Bengal Fan, which covers an area of 3000 x 1000 km to a maximum thickness of 16.5 km (Curray et al., 2002) and has been estimated to hold about 10–20% of the total terrestrial organic carbon sequestered in global marine sediments (France-Lanord and Derry, 1997; Galy et al., 2007). In fan sediments, lighter and darker organic-rich bands reflect biomass burial associated with high-discharge events, which support high OC burial efficiency by delivering recalcitrant biomass swiftly through rivers and then burying it under inorganic sediments (Raymond et al., 2016; Battin et al., 2023). The circulation of the Bay Bengal region also creates a zone of low O_2 in the water column, from ~50–200 m depth (Rao et al., 1994).

In the Bengal Fan and other major river systems, much of the terrestrial OC is transported to the deep sea by episodic turbidity currents. Individual events are capable of transporting massive amounts of material thousands of kilometers down submarine canyons, where rapid burial can lead to efficient (60-100%) organic carbon preservation (Masson et al., 2010; Baker et al., 2024). In the Congo River system, which carries roughly 7% of all global riverine organic carbon, turbidity currents eroded ~2.7 km^3 of sediment over a single year, which is equal to ~25% of the global annual suspended sediment river flux

https://doi.org/10.5194/egusphere-2025-6086
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(Talling et al., 2022; Hage et al., 2024). Both river deltas and adjacent canyon systems can thus be significant repositories for organic carbon.

In principle, rapidly accumulating sediments could support carbon storage for CDR at a continental scale. Large sediment transport systems exhibit physical dynamics and mineral interactions that add complexity to carbon quantification and the prediction of environmental impacts, necessitating additional research into these processes. River deltas and canyons are also less isolated from the open ocean than anoxic basins, which offers advantages for carbon transport and storage scale but also presents challenges for governance, carbon quantification, and monitoring. Consequently, although sediment burial represents a potentially viable pathway for MACS development in the Bay of Bengal region and elsewhere, it is not the focus of the scaling analysis below.

Additional sites may be valuable scientific resources and analogs for the development of MACS datasets and MRV systems. Fjords with anoxic bottom water, for example in Norway, British Columbia, and Alaska, provide insights into anoxic organic carbon burial over centuries to millennia.

3. Key Criteria to Define the Maximum Scale of MACS for CDR

The goal of this section is to evaluate mechanisms that could limit the potential global scalability of MACS for CDR. Each of the following sections covers one of five criteria that could define such a maximum deployment scale: (A) Durable storage site capacity (physical considerations); (B) Biomass sources and logistics; (C) Oxygen and sulfide impacts at the redoxcline; (D) Greenhouse gas balance; and (E) Impacts on dissolved organics or nutrients in the oxic zone. For each criterion, we discuss the fundamental processes involved, potential risk mechanisms associated with MACS, the current state of knowledge, and key knowledge gaps. In this analysis, we focus on quantifiable issues of mass, energy, and economics to identify first-order constraints on the potential scale of MACS. These results are one component of societal discussions related to the social, technical, political, ecological, and cultural aspects of decision making around climate interventions (Cooley et al., 2023; Nawaz et al., 2023; American Geophysical Union, 2024).

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

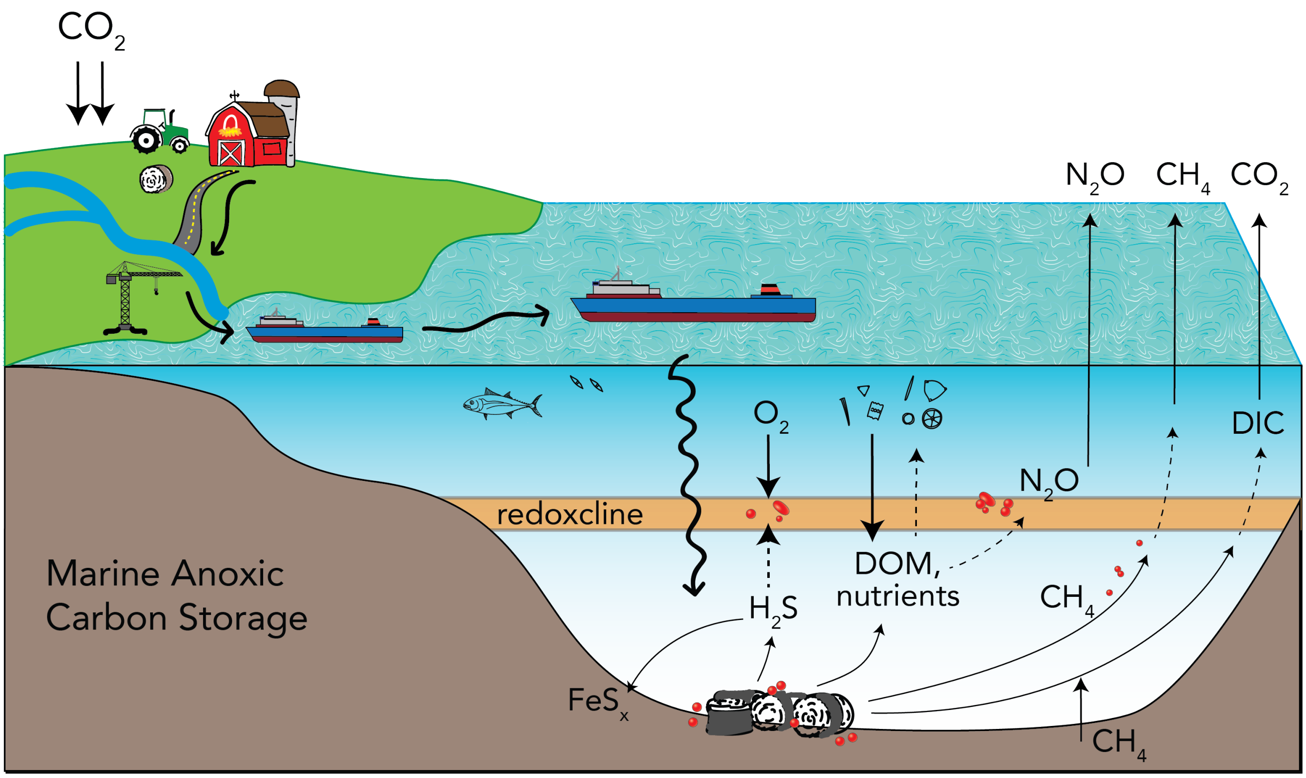

Figure 1: Important processes related to marine anoxic carbon storage. Diagram is not to scale and is generalizable to both the Black Sea and brine pool systems. Red spots indicate important microbial processes discussed below. Dashed lines represent limited mixing across stratified water masses.

3.1 Durable storage site capacity

Estimating the physical capacity of the Black Sea and brines, as well as other sites in the future, requires quantifying the total volume of water masses that have favorable characteristics for carbon sequestration. Although many factors will contribute to an evaluation of site favorability, one essential criterion is durability, which describes the amount of time that sequestered carbon will be isolated from the atmosphere. Some of the same factors that affect durability – like mixing, stratification, and circulation – also control the timescale and intensity of potential environmental impacts in the oxic ocean. The physical processes limiting the mixing of deep, anoxic water at potential sites are therefore foundational controls on global MACS scale.

Critically, to predict how long the products of biomass breakdown will take to reach the oxic ocean or atmosphere (i.e., their durability), we must understand how deep waters are likely to circulate and mix in the future. This requires more than simply knowing how long today's deep waters have been isolated from the surface. In the Black Sea, changes in surface temperatures, hydroclimate and river inputs, wind forcings, and sea level may affect stratification and circulation of the storage environment over decades to centuries (Capet et al., 2013, 2016; Ilicak et al., in prep). Secondary effects on circulation and durability could be caused by changes in physical dynamics or turbulence on the seafloor due to physical structures or modified

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

bathymetry. Heat generated during microbial breakdown could in theory influence deep mixing, but this effect is likely small (Raven et al., 2024). Additional considerations for defining favorable regions of the deep Black Sea are sediment characteristics, sedimentation rate, subseafloor processes (e.g. mud volcanism and gas seepage), and the avoidance of sensitive locations.

Risk mechanisms

- Changes in basin circulation (more frequent upwelling and/or faster vertical mixing) could transport breakdown products to the redoxcline and atmosphere.
- Added biomass could fill or physically displace storage volume.

Current level of understanding

The seawater or brine volumes of the reservoirs discussed here (Black Sea, Orca Basin) are well known. The abyssal Black Sea as a whole is not physically storage-limited at the Gt scale (Murray et al., 1991; Raven et al., 2024), but sub-regions of the Black Sea with particularly favorable conditions for storage remain to be quantified.

In contrast, for brines, physical volumes may be important limits to the potential scale of biomass storage. The MACS capacity for the Orca Basin may be roughly 0.1 to 0.5 Gt CO_2e (Martinez et al., in prep). A comprehensive global survey of potentially favorable brine sites for CDR is not yet available, and estimates are largely dependent on the availability of quality data at the appropriate scales for the system. Additional work is needed to expand and improve bathymetric, hydrographic, and circulation datasets and to apply modern data science tools to extract as much information from those data as possible.

Key unknowns

1. Over what timescales are anoxic water masses isolated from the surface? How will the isolation of these water masses change in the future?
2. What is the global volume of deep hypersaline brine pools? Which brine types (chemistry, temperature) are most conducive to MACS?

3.2 Biomass Sources and Logistics

The main biomass sources considered here are terrestrial plants, including: (1) agricultural materials/wastes that can be sustainably removed from fields and processing facilities; and (2) excess woody material produced during ongoing forest management or wood processing activities. In principle, MACS could utilize a wide diversity of biomass materials (e.g., the biogenic portion of municipal solid waste streams, nuisance algae like Sargassum blooms), although an acceptable level of

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

resistance to breakdown (recalcitrance) and density/specific gravity would need to be demonstrated, as well as an absence of contaminants.

The quantity of available biomass sources will depend on several management factors (whether agronomic, forest, or waste management), all of which vary in a geospatial context. In addition to knowing the net annual production of a particular biomass residue, it will be necessary to determine what proportion of that material could be removed from land without detrimental effects on soil health and nutrient budgets, and to consider the energy and infrastructure impacts of the full life-cycle assessment. Here, "sustainable" biomass sources with acceptable lifecycle CO_2 emissions profiles would define maximum reasonable sources.

Additionally, the practical availability of biomass resources will depend on socioeconomic factors, including consideration of their competing uses. This analysis requires some socioeconomic projections and assumptions, but energy-efficient supply chains for MACS are a priority because they would be needed across many potential future scenarios. Practical and efficient biomass transportation networks are likely to be a key determinant of whether any particular biomass source is favorable for MACS.

Greenhouse gas emissions associated with biomass transportation and handling are also discussed under criterion D, below.

Risk mechanisms

- Excess biomass removal from agriculture could lead to soil degradation, soil organic carbon loss, biodiversity loss, and/or potential loss of future production.
- Biomass use for CDR could modify incentives for land use and economic development. Changes driven by those incentives could be deemed negative for a diversity of political, economic, and social reasons and represent a limit to scale.
- Competition with alternative uses for lignocellulosic biomass could limit the quantity available for CDR.
- Resource-intensive transportation and handling of biomass could undermine carbon removal efficiency or cost-effectiveness.
- Biomass transportation activities could lead to undesirable side effects like truck traffic, port crowding, or ship noise, impacting sea life around ports or deployment locations.

Current level of understanding

Substantial work has been done to understand availability and suitability of biomass resources for other purposes like biofuels, which provides a largely complete conceptual framework, particularly in the United States and Europe (Pett-Ridge et al., 2023; Huntington et al., 2024). However, other aspects are more specific to ocean-based CDR and remain

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

underdeveloped, most notably a geospatial context that considers the proximity of the potential storage area, transportation distances, water-based logistics, navigation safety, and the total carbon footprint of carbon removal activity. This necessitates place-based consideration of alternative biomass uses.

Existing and potential future sustainable biomass resources are relatively well characterized for the purposes of transportation logistics analysis. Price, type, and quantity of biomass at a county-level resolution is available in the United States (Langholtz, 2024), and similar datasets exist for the European Union (Rosa et al., 2021). Datasets for other regions are more scarce and may require additional evaluation.

Analyses of regional and continental-scale biomass supply and logistics analyses have been performed at increasingly realistic levels of detail in recent years, especially in the United States. Specifically, researchers can leverage advanced geospatial data related to biomass availability, available transportation modes, and potential locations for collection, aggregation, and processing to propose transportation logistics that meet varying economic and social criteria (optimizing for cost, carbon emissions, truck traffic, etc.) (Pett-Ridge et al., 2023; Huntington et al., 2024; Oyedeji et al., 2025). Modeling can typically be completed using linear or mixed-integer linear programming, reducing computational intensity.

Biomass sources may represent a practical limit on the CDR potential for MACS, depending on actual environmental, political, and economic trajectories. For "maximum theoretical" values, numbers as high as 10 Gt CO_2e/yr have been proposed (Zeng et al., 2024). In the context of actual supply chains, competing biomass uses, and regulatory structures, however, additional work is needed to quantify potential biomass sources for alternative deployment scenarios.

Key unknowns

3. Where could effective supply chains link biomass source locations, waterways that minimize the interference with shipping and other marine uses, and anoxic basin storage sites?
4. Which potential biomass sources and transportation alternatives could provide the highest efficiency in terms of net carbon removal? How much of the terrestrial biomass that is produced today would meet a minimum threshold for carbon efficiency via MACS?
5. How do potential biomass sources and transportation options differ in terms of their energy requirements, infrastructure needs, and projected economic impacts? What other social, political, and environmental factors will affect the feasibility and favorability of potential biomass sources?

3.3 Greenhouse gas balance

Greenhouse gas budgets relevant to MACS include CO_2, methane (CH_4), and nitrous oxide (N_2O). Methane and nitrous oxide are high priorities for mitigation because their global warming potentials exceed that of CO_2 by 25x and 273x, respectively. Greenhouse gases can be released as part of the biomass sourcing process as well as during biomass breakdown.

Additionally, both the Black Sea and Orca Basin are naturally methane-rich environments (Wiesenburg et al., 1985; Reeburgh et al., 1991), and changes to natural rates of methane release would contribute to the net greenhouse gas impacts of a hypothetical MACS project.

Transportation and, optionally, drying and densification of biomass are responsible for CO_2 and potentially methane emissions. Transportation is typically one of the largest sources of greenhouse gas emissions in envisioned biomass carbon removal and storage processes, although biomass processing can be significant depending on biomass type, handling process, and the efficiency of local electricity. Greenhouse gas emissions associated with supply chains are reflected in calculations of carbon efficiency (see criterion B).

CO_2 is one component of the dissolved inorganic carbon (DIC) pool in seawater, along with bicarbonate (HCO_3^-) and carbonate ion (CO_3^{2-}). The equilibration among these species is a function of pH and alkalinity, and only CO_2 exchanges with the atmosphere (Wanninkhof et al., 1999). Changes in these parameters will therefore affect the capacity of surface waters to store carbon, as has been observed over storm and seasonal cycles in the Black Sea (Müller et al., 2016; Voynova et al., in prep). DIC is produced by a wide diversity of anaerobic microbial metabolisms through the oxidation of organic C (Middelburg et al., 2020). In the deep Black Sea, DIC concentrations of approximately 4.1 mM reflect accumulation over the long residence time of deep waters (Hiscock and Millero, 2006). Orca Basin DIC concentrations also exceed 4.5 mM due to long-term trapping of biomass breakdown products.

Methanogenesis is a fermentative metabolism that can dominate in environments without electron acceptors like sulfate. Even at sulfate-replete sites, some groups of methanogens (e.g., methylotrophs) can generate methane by taking advantage of “non-competitive” substrates like the methoxy groups in lignins. Importantly, however, only a tiny fraction of the methane that is released from the seafloor reaches the atmosphere because methane can be readily oxidized by methanotrophic microorganisms under both anoxic and oxic conditions (Hinrichs and Boetius, 2002). Naturally-occurring populations of methanotrophs in both the Gulf of Mexico and the Black Sea appear to have the capacity to respond rapidly to sudden influxes of methane and prevent the vast majority from reaching the atmosphere (Kessler et al., 2011; Schmale et al., 2011). Under some conditions, methane can also be released from vents on the seafloor as bubbles. Methane bubbles largely dissolve into the surrounding water as they rise, and if released at high pressure, mass transfer calculations suggest that such dissolution increases (Haeckel et al., 2015). As such, ebullition is not expected to release significant amounts of methane into the atmosphere if the water column is > 100 m deep (McGinnis et al., 2006; Grilli et al., 2021).

Despite its active internal methane cycle, net fluxes of methane from the Black Sea to the atmosphere, excluding coasts, is small: ~0.05 Tg CH_4/yr (Reeburgh et al., 1991; Kessler et al., 2006). Methane in deep Black Sea water (~11 µM) derives from both microbial and thermogenic sources, including hydrocarbon seeps and mud volcanoes caused by the heating and diagenesis of deeply buried sediments (Starostenko et al., 2010). Microbial methanogenesis in the sediments is an important source of methane to the basin above ~1500 m depth, while sediments in the deepest part of the basin are methane sinks (Reeburgh et al., 1991). Despite large and sometimes sudden inputs of methane to the water column from seeps and tectonic activity, the residence time of methane in the deep Black Sea is around 75 years due to efficient reoxidation (Kessler et al., 2006; Schmale et al., 2011). For MACS at deep sites, this suggests that CO_2 rather than methane is the greenhouse gas with the largest potential to be released to the atmosphere.

In Orca Basin, methane concentrations are more than 600 µM in the brine but drop sharply to less than 10 nM at the redoxcline (Wiesenburg et al., 1985). Methane is effectively trapped by the density interface and may also be oxidized within the redoxcline. The vast majority of this trapped methane likely derives primarily from active hydrocarbon seepage of biogenic gas into the brine pool (Sackett et al., 1979). Although sulfate is abundant in both the brine and pore water, methylotrophic methanogens are present and appear to contribute methane at low rates (Zhuang et al., 2016; Nigro et al., 2020). Nevertheless, in the absence of physical perturbations to the density interface, the addition of methane and CO_2 to the brine from MACS is unlikely to significantly change fluxes of these greenhouse gases to the overlying water column.

Nitrous oxide can be produced by multiple N-cycling reactions near oxic-anoxic interfaces, including ammonium oxidation (nitrification) and nitrate/nitrite reduction (denitrification). In the upper parts of the redoxcline where O_2 concentrations are low, chemoautotrophic microorganisms oxidize upward-diffusing ammonium with O_2 to produce N_2O along with nitrite and nitrate (Kim and Craig, 1990). In the deeper parts of the redoxcline, upward-diffusing DOC meeting downward-diffusing nitrate from overlying oxic waters may also result in the production of N_2O by incomplete denitrification. However, denitrification can also remove N_2O by reducing it to N_2 (Cohen and Gordon, 1978; Haas, 2020). All of these processes have been shown to occur in different anoxic water columns in different combinations (Westley et al., 2006; Wenk et al., 2016; Haas, 2020). Notably, in the Black Sea, the net production of N_2O is roughly equivalent to open-ocean sites; nearly all of the N_2O generated from nitrification is consumed by denitrification (Westley et al., 2006). Any impact of MACS on N_2O cycling would likely depend on the degree to which biomass breakdown increases ammonium and organic carbon fluxes to the redoxcline, as modulated by the physical processes that influence redoxcline structure. Impacts of MACS on the nutrient and DOM cycles are also discussed under criterion E.

Risk mechanisms

- DIC from biomass breakdown that mixes to the surface could equilibrate with the atmosphere and be released in part as CO_2.
- Methanogenesis could release CH_4 that mixes to the surface before it is oxidized, either aerobically or anaerobically, by microbes.
- Methane in sedimentary reservoirs such as gas hydrates or brines could be disrupted by physical deposition of biomass materials or monitoring activities. This methane could impact the atmosphere if it were transported to the surface faster than it could be oxidized.
- The net production of N_2O could increase due to enhanced fluxes of ammonium (fueling nitrification) or DOM (driving denitrification) to the redoxcline.

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

Current level of understanding

Greenhouse gas emissions associated with biomass supply chains have been the subject of intensive work by others, and procedures for their calculation are well established (e.g., Puro.earth, 2025). Mapping emissions onto specific biomass source scenarios for MACS remains an area for future work (criterion B).

Rates of greenhouse gas production by microbial processes have been studied in many natural environments, but a hypothetical MACS deployment would fall outside of the range of most natural analogs. Experimental and/or observational data are needed to quantify at least two key parameters – breakdown rate and the recalcitrant fraction of carbon (Westrich and Berner, 1984; Arndt et al., 2013) – for conditions of sudden biomass addition to Black Sea or brine environments.

The physics of gas exchange at the sea surface are well understood, but local variations in temperature, productivity, circulation, and other factors lead to complex dynamics. For example, in the summer and during most of the warm season, and increasingly under a warming climate, shelf seas like the Black Sea can be classified as weak net CO_2 sources to the atmosphere, except during the spring-early summer productive seasons, when surface waters can become CO_2 sinks. Over decadal timescales, increases in dissolved inorganic carbon are observed in the oxic layers (Voynova et al., in prep). Determining the limits of the processes driving DIC accumulation and their effects on the layers below the redoxcline will be essential to estimate future CO_2 fluxes from the Black Sea. Similarly, given the multitude of processes impacting N_2O production and consumption at redox interfaces, coupled physical-biogeochemical models will be needed to understand how the effects of changing ammonium and DOM fluxes are modulated or overprinted by mixing and climate dynamics.

Another important area for future evaluation considers non-steady-state processes in the natural methane cycle. The "baseline" conditions for both environments are subject to natural disruption from tectonics (earthquakes, venting, diapirism), sedimentary processes (landsliding), and circulation change (deep water ventilation). These intermittent events present a challenge for future predictions of the "natural" methane cycle and the attribution of any direct or indirect impacts from MACS.

Key unknowns

6. At what rate will DIC and methane be produced from biomass at the seafloor? What fraction of biomass is effectively resistant to breakdown under site conditions?
7. What is the efficiency of methane oxidation in response to a hypothetical increase of methane emissions from MACS?
8. How susceptible are seafloor methane hydrates and brine interfaces to physical disruption? How is excess methane mixed and transported in the context of non-steady-state processes?

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

3.4 Oxygen and sulfide impacts at the redoxcline

The availability of dissolved O_2 is a fundamental control on most marine ecosystems and plays a key role in the composition and function of marine communities. Many marine animals exposed to hypoxic waters (i.e., at concentrations below ~60 µM) experience physiological stress, although recent studies have identified dense aggregations of cold-water corals in hypoxic waters (Hebbeln et al., 2020). Oxic respiration is severely curtailed, even for microbes, below ~5 µM (0.2 mg/L, 0.14 mL/L) (Vaquer-Sunyer and Duarte, 2008). Anoxia is often practically defined by the detection limit of standard sensors and titrations at 1–2 µM, although microbial O_2 dynamics remain significant into the range of nanomolar O_2 concentrations (Tiano et al., 2014; Deutsch et al., 2024). Below even these trace concentrations, fully anoxic water comes in multiple flavors. In the deep Black Sea and many other anoxic basins, seawater is sulfidic, having accumulated dissolved H_2S from microbial sulfate reduction. In certain sites like Orca Basin, anoxic seawater is ferruginous, containing high concentrations of dissolved Fe^{2+}, but no detectable O_2 or H_2S. In either case, the interface between the well-oxygenated surface layer and an underlying anoxic zone, the "redoxcline" (Fig. 1), is a critical boundary where largely microbial processes exchange electrons between reduced species like sulfide and oxidized species like O_2. This O_2 gradient is often co-located with a density gradient (pycnocline) that impedes mixing, especially for brines.

Biomass breakdown by sulfate reduction below the redoxcline produces H_2S, which has several possible fates in the environment. In the presence of dissolved iron or poorly crystalline iron (oxy)hydroxide minerals, which are common in sediments, sulfide can react to form iron monosulfide or disulfide minerals, which tend to precipitate as solids. Alternatively, sulfide that is transported toward the redoxcline can be oxidized by a variety of biological and abiotic mechanisms, typically to sulfate (Zhang and Millero, 1993). Sulfide may also react with functionalized organic molecules and trace metals, modifying their chemical properties (Dyrssen and Kremling, 1990; Saager et al., 1993). If it invades otherwise oxygenated areas, dissolved sulfide is toxic to many marine organisms at micromolar concentrations. In Eastern Boundary upwelling systems with anoxic waters, for example, episodic sulfide plumes from sediments are implicated in fish population losses (Schunck et al., 2013).

In the Black Sea, processes at the shallow (~150 m) redoxcline are particularly important because they determine whether chemical species produced in the anoxic zone reach the ecologically rich and economically valuable oxic layers above. The position of the oxic-anoxic interface in the water column has varied over time in response to changes in mesoscale activity (Zarokanellos, in prep.), agricultural runoff (eutrophication) and hydroclimate (Konovalov and Murray, 2001; Pakhomova et al., 2014). Regardless of mechanism, elevated concentrations of sulfide in the upper anoxic water column would be expected to cause greater O_2 loss and a shoaling of the redoxcline, compressing this upper oxic layer and influencing biodiversity and nutrient cycles (Gilly et al., 2013; Köhn et al., 2022), as appears to have occurred throughout the Holocene (Huang et al., 2000). Understanding these mechanisms is essential for estimating how MACS-related perturbations could interact with the existing anthropogenic pressure in the Black Sea.

The Orca Basin, in contrast, is ferruginous, with concentrations of dissolved Fe^{2+} as high as 26 µM (Wiesenburg et al., 1985). Sulfide produced by sulfate-reducing microorganisms, either within the redoxcline or near the sediment-water interface, reacts rapidly with this dissolved iron to precipitate FeS solids, which accumulate to remarkable concentrations in sediments (>1 wt%; Hurtgen et al., 1999). The position of the Orca Basin redoxcline is primarily driven by salinity, but large changes in sulfide fluxes could affect the chemistry of the brine and the distribution of microbial communities within its chemical gradients.

Risk mechanisms

- Sulfide released at depth and transported to the redoxcline could lead to the consumption of O_2 or nitrate at this boundary. Expansion of the anoxic zone could drive loss of habitat for benthic or pelagic communities.
- Localized sources of sulfide could drive acute ecosystem impacts through episodic mixing events.

Current level of understanding

In general, biomass breakdown rates are sensitive to the properties of both the biomass and the environment (Arndt et al., 2013). Rates of sulfide production during biomass breakdown have been studied in a wide range of anoxic environments, although only a subset of these studies consider the combination of terrestrial materials and anoxic marine conditions (Keil et al., 2010; Blair and Aller, 2012; Hage et al., 2020). Additional biomass breakdown experiments are needed to constrain breakdown rates and the recalcitrant fraction for potential MACS materials.

There is a solid foundation of academic literature available to estimate some important secondary reactions between sulfide and iron species, trace metals, organics, and oxidants. However, several potentially important chemical reactions remain difficult to predict in the environment. For example, both iron oxide and iron sulfide minerals can contribute to organic carbon preservation in sediments because organic molecules within and on the surfaces of these minerals may be protected from microbial breakdown (Lalonde et al., 2012; Nabeh et al., 2022). Changes in redoxcline position and sulfide availability can dissolve and re-form those minerals, but the net effect of this reaction on carbon storage has not yet been quantified.

Another area of active research focuses on the interactions between O_2, sulfide, and other redox-active species at the redoxcline, in the context of dynamic physical mixing and transport processes that span the timescales of eddies, upwelling plumes, or decadal climate oscillations (e.g., Pakhomova et al., 2014; Buchanan et al., 2023). In the Black Sea, observations of the relationships between oxygen depletion and sulfide dynamics near the redoxcline of the Black Sea illustrate the impacts of both natural variability and human activities on this system (Konovalov et al., 2006). Twentieth-century eutrophication in particular serves as a valuable case study for the response of the Black Sea to changes in organic matter availability and sulfide production (Kroiss et al., 2006; Capet et al., 2013), although the analog is imperfect as the enhanced rates of sulfide production from eutrophication occur largely in the water column, while those from hypothetical MACS interventions would occur on the deep seafloor.

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

Key unknowns

9. What proportion of excess terrestrial biomass added to an anoxic environment will be consumed through microbial sulfate reduction over short (days to months) and long (years+) timescales?
10. What are the sinks for sulfide in the deep Black Sea and seafloor brines? How do changes in sulfide flux affect net carbon storage?
11. How would ecosystems within and overlying the Black Sea redoxcline respond to changes in deep sulfide flux?

3.5 Impacts on nutrients or dissolved organic matter (DOM) in the oxic zone

A hypothetical MACS deployment could impact nutrients and/or DOM in the surface ocean at two different points in the deployment process. Nutrients could leach rapidly during initial wetting and placement of terrestrial biomass, or nutrients released at the seafloor could be transported to the surface by upwelling water and mixing (Raven et al., 2025). The specific nutrients and organic molecules released from biomass during both stages will depend on the biomass type and its preparation.

Concentrations of most nutrients in lignocellulosic biomass are low relative to marine organisms, with C:N ratios in wood often ≥300:1 (Cowling and Merrill, 1951). Nevertheless, at scale, the breakdown of this biomass could impact marine ecosystems through changes in nutrient availability. Essential nutrients for primary production in the photic zone include the macronutrients nitrogen (N) and phosphorus (P) as well as silica (Si), iron (Fe), and other trace metals. Nutrients are primary controls on the amount and distribution of net primary productivity (NPP) in the surface ocean, and the relative availability of different nutrients in the environment affects ecosystem structure. Changes to either the amount or the type of NPP (e.g., its chemical composition, seasonality, or export ratio) have direct ramifications for higher trophic levels like marine mammals and fish (Berger et al., 2023; Tagliabue et al., 2023).

The nitrogen cycle is also an active participant in redox cycles at the redoxcline, influencing both N_2O production (criterion C) and redoxcline position due to the consumption of O_2 by NH_4 oxidation (criterion D). Many micronutrients are also actively cycled near the redoxcline, including iron and many trace metals. These reactions will determine if and how changes in nutrient concentrations in deep water translate into discernable changes in oxic environments.

Ecosystems both within and below the photic zone are highly sensitive to the availability of dissolved organic matter (DOM). DOM can be released from essentially all forms of biomass through senescence, leaching, or heterotrophic activity in seawater, although lignocellulosic biomass is generally less soluble in seawater than biomass from marine primary producers. Many anthropogenic organic contaminants also fall in the category of DOM, including many agricultural chemicals like pesticides as well as products of the breakdown of (micro)plastics. From the perspective of marine organisms, the vast pool of molecules categorized as DOM includes energy sources (food), building blocks for biosynthesis (e.g., amino acids), and key reactants in cellular processes (e.g., B vitamins, signalling molecules). Additionally, some DOM molecules are acidic and can cause pH change (e.g., carboxylic acids), while others can complex and stabilize trace metals. Therefore, it will be important

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

to understand both the amount and the chemical properties of any DOM added to the environment in order to predict ecosystem response.

In the Black Sea, concentrations and ratios of nutrients and DOM in the photic zone have experienced dramatic change since at least 1960, in parallel with major ecosystem shifts. The primary source of nutrients to photic zone ecosystems is mixing with deep water (Ludwig et al., 2009). Products of biomass breakdown have accumulated in deep Black Sea water due to its slow turnover time, reaching up to 100 µM ammonium and 8 µM phosphate (Brewer and Murray, 1973; Konovalov et al., 2006). In the seasonal Cold Intermediate Layer that feeds much of the surface nutrient budget, concentrations of inorganic N roughly tripled between the 1950s and 1980s and then recovered to levels around 2x pre-eutrophication values (Mikaelyan et al., 2013; Pakhomova et al., 2014). These dramatic changes in nutrient availability have been associated with changes in both the amount and the source of primary production in the basin (dinoflagellates, diatoms, and/or coccolithophorids; Mikaelyan et al., 2013). Changes in the chemical composition, particle export, and seasonality of primary producers all have downstream effects on ecosystems, as seen in the Black Sea in recent decades (Oguz and Gilbert, 2007).

Orca Basin contains even more highly concentrated levels of biomass breakdown products than the Black Sea. The deep brine contains ≥450 µM dissolved OC, ≥520 µM ammonium, and ≥50 µM phosphate (Van Cappellen et al., 1998; Martinez et al., in prep), as nutrients and DOM released within the brine are effectively trapped within that water mass without being consumed or metabolized. In contrast, the overlying water column is generally oligotrophic, although surface nutrients and primary productivity vary in response to seasonal cycles and mesoscale eddies (Damien et al., 2021).

Risk mechanisms

- Changes in nutrient levels or ratios could favor different primary producers, drive blooms of undesirable algae, displace species, or generate excess production that could impact O_2.
- Biomass-derived DOM that mixes to the surface could stimulate or alter ecosystem processes in the oxic zone.
- Biomass could introduce contaminants such as plastics, pesticides, herbicides or fungicides that could affect marine ecosystems.

Current level of understanding

Macronutrient–NPP relationships are relatively well implemented in marine ecosystem models. Basic biogeochemical and redox budgets have been created for the Black Sea (Brewer and Murray, 1973; Konovalov et al., 2006), and relatively large datasets exist to explore the relationships between twentieth-century fertilization trends and ecosystem impact (e.g., Pakhomova et al., 2014; Stanev et al., 2022). Strategically placed coastal stations in regions with large freshwater influxes, like the western Black Sea, have been helpful to observe the impacts of seasonal and interannual changes in river

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Preprint. Discussion started: 18 December 2025

discharge, and of coastal upwelling and extreme events on coastal ecosystems (Voynova et al., in prep). Modeling tools are also poised to explore potential impacts of hypothetical MACS deployments in the context of mesoscale and dynamic circulation patterns.

For brines like Orca Basin, water mass trapping at the redoxcline limits the potential impacts of nutrient and DOM release on the overlying environment from long-term biomass breakdown. The primary mechanism for impact on the oxic zone from MACS is more likely to be short-term effects during initial wetting and placement. Short-team leachates from some biomass types have been characterized, but potential ecosystem sensitivities to various large-scale deployment strategies still require evaluation (Martinez et al., in prep). Nutrient and DOM impacts may be minimized through engineering choices and biomass selection for low nutrient content in leachate.

In the context of large-scale deployment, several additional mechanisms for nutrient and DOM impacts could become important and would require additional research attention. DOM from specific terrestrial sources may differ from pre-existing DOM in ways that may impact the availability of essential trace metals. Certain DOM molecules may also either stimulate or repress anaerobic secondary production (heterotrophic growth). And finally, the release of organic acids could also have local effects on pH and alkalinity, which also has implications for measurement and monitoring methods that rely on those parameters.

Key unknowns

12. Which nutrients and organic molecules are released from biomass during deposition? Which biomass alternatives or biomass processing measures minimize release?
13. At what rate would DOM and macronutrients be released during biomass breakdown on the seafloor?
14. How reactive is the released DOM in the environment? Is it labile for anaerobic heterotrophs? How does it affect the speciation of trace metals?

	Criterion	**Risk Mechanisms: Potential Limit to MACS Scale**	**Relevance/Priority**	
			Black Sea	**Brines**
A.	**Durable storage site capacity**	Changes in basin circulation (more frequent upwelling and/or faster vertical mixing) could enhance the transport of breakdown products to the redoxcline and atmosphere.	high	low
		Added biomass could fill or physically displace storage volume.	low	high
B.	**Biomass source availability**	Excess biomass removal from agriculture could lead to soil degradation.	low	low
		Biomass use for CDR could modify incentives for land use and economic development.	low	low
		Competition with alternative uses for lignocellulosic biomass could limit the quantity available for CDR.	mid	mid
		Resource-intensive transportation and handling of biomass could undermine carbon removal efficiency or cost-effectiveness.	high	high

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

		Biomass transportation activities could lead to undesirable side effects like truck traffic, port crowding, or ship noise, impacting sea life around ports or deployment locations.	mid	mid
C.	**Greenhouse gas balance**	DIC from biomass breakdown that mixes to the surface could equilibrate with the atmosphere and be released in part as CO_2.	high	low
		Methanogenesis could release CH_4 that mixes to the surface before it is oxidized, either aerobically or anaerobically, by microbes.	mid	mid
		Methane in sedimentary reservoirs such as gas hydrates or brines could be disrupted by physical deposition of biomass materials or monitoring activities, which could impact the atmosphere if it reached the surface faster than it could be oxidized.	high	high
		The net production of N_2O could increase due to enhanced fluxes of ammonium or DOM to the redoxcline.	low	low
D.	**Oxygen and sulfide impacts at the redoxcline**	Sulfide released at depth and transported to the redoxcline could lead to chemical change at this boundary. Expansion of the anoxic zone could drive loss of habitat for benthic or pelagic communities.	high	mid
		Localized sources of sulfide could drive acute ecosystem impacts through episodic mixing events.	low	low
E.	**Nutrient or DOM impacts in the oxic zone**	Changes in nutrient levels or ratios could favor different primary producers, drive blooms of undesirable algae, displace species, or generate excess production.	mid	low
		Biomass-derived DOM that mixes to the surface could stimulate or suppress ecosystem processes.	high	low
		Biomass could introduce contaminants such as plastics, pesticides, herbicides or fungicides that could affect marine ecosystems.	mid	mid

Table 2: Summary of risk mechanisms that could represent limits to the global scale of MACS for CDR. Columns at right indicate mechanisms that are expected to have low (white), moderate (grey) or high (black) relevance for determining potential scale at each site.

4 Summary and Next Steps

Unchecked climate change is a clear threat to marine ecosystems as well as the foundational activities on which human societies rely (Doney, 2010; Müller et al., 2011). In the absence of net-negative CDR, there is a meaningful risk that global climate exceeds critical and effectively irreversible planetary thresholds with devastating consequences to human societies (Armstrong McKay et al., 2022; Drijfhout et al., 2025). Therefore, globally, we must find a way to remove hundreds of gigatonnes of CO_2 from the atmosphere in the coming decades in addition to achieving immediate and dramatic emissions reductions (IPCC AR6, 2021). To chart the most effective and ethical path forward, it will be essential to transparently balance the risks of different potential climate inventions against the risks of inaction.

Marine Anoxic Carbon Storage (MACS) may have the potential to contribute substantially to global CO_2 removal goals by sequestering biomass in restricted, anoxic marine reservoirs. The feasibility and societal acceptability of this pathway will depend on many factors, but there are at least five fundamental constraints on scale that will need to be met by any proposed actions. These include: sufficient durable (≥kyr) storage capacity, suitable and sufficient biomass sources, an

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

acceptable greenhouse gas balance, limited impacts on oxygenation, and acceptable impacts on nutrients and DOM in the oxic zone. For each of these criteria, there are remaining uncertainties that need to be addressed in order to estimate the potential responsible global scale of MACS for CDR. Broadly speaking, we have a basic understanding of many of the key processes in this system as they operate today and in the recent past, but we still need to develop models and analogs to answer questions concerning the outcomes of much larger-scale biomass addition for CDR.

Based on what we know so far, it may be possible for MACS to exceed the 0.1–1.0 Gt/yr CO_2e significance thresholds for moderate or high scalability (NASEM, 2022). To increase confidence in this assessment, we need to answer several top-priority questions that have particularly high leverage in this calculation:

- How is the circulation of the Black Sea likely to change in a changing climate?
- What is the global volume of seafloor brines with favorable properties for MACS?
- How do the carbon and energy efficiency of potential biomass sources compare, and how much biomass is well suited for MACS?
- What is the long-term (kyr) preservation efficiency of lignocellulosic biomass for various deployment scenarios?
- What is the potential for physical disruption of sediments and/or methane reservoirs, and how can it be minimized?
- How would biomass breakdown products impact biogeochemical processes and the carbon system at the redoxcline, in the context of its physical dynamics?
- What is the potential for MACS in anoxic sedimentary systems like river deltas and submarine canyons?

These key questions represent the mission ahead for collaborative research across the field and are urgent areas to prioritize in future research. They have major leverage on the calculation of potential MACS scale because they may identify early "no-go" outcomes, or limits on feasibility. This list is not intended to be exhaustive of the scientific knowledge gaps worthy of detailed attention, especially related to the response and ecosystem function of deep microbial communities. Instead, these questions interrogate to what extent such detailed scientific efforts would potentially unlock real and positive benefits to climate and society.

In parallel with a scientific effort to address these unknowns, broad and interdisciplinary work is essential to understand the priorities and interests of local communities that would be impacted by MACS activities at scale. Building on the global-scale risk assessment initiated here, robust protocols and metrics will need to adapt to those local considerations while ensuring that environmental protection standards are met regardless of political jurisdiction. Multilateral international collaboration among coastal nations will be essential for effective MACS implementation, as hydrology and circulation patterns do not recognize national boundaries.

We have focused on anoxic basins in this first analysis, which have several key practical advantages due to their relative isolation from the global ocean. Nevertheless, enhanced terrestrial OC burial events in Earth history have been centered

https://doi.org/10.5194/egusphere-2025-6086
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primarily in large coastal river and submarine canyon systems, motivating future research in these areas. For example, in the Bay of Bengal region, joint research programs involving India, Bangladesh, and Myanmar could help characterize biogeochemical processes governing carbon cycling in the region and facilitate the development of a unified regional framework for MACS deployment that accounts for shared hydrological connectivity, equitable benefit distribution, and coordinated monitoring protocols.

MACS at the scale of 1 Gt CO_2e/yr would represent a roughly three-fold acceleration of excess OC burial relative to 40,000-year averages for the PETM recovery (Papadomanolaki et al., 2022) or a 30% increase in modern global OC burial rates in the ocean (Dunne et al., 2007; Table 1). These are large changes, but they fall within the apparent range of Earth's natural variability. Geologically recent historical analogs like the PETM (~56 Mya) provide opportunities to understand how large-scale terrestrial OM burial events impact climate, ecosystems, and other feedbacks in the Earth system, and they provide confidence that large-scale terrestrial OM burial in the ocean can effectively lower atmospheric CO_2.

Important scientific questions need to be answered before we can make informed decisions about if and how to deploy MACS at scale for CDR. Nevertheless, there is a clear path toward addressing those unknowns, and substantial research progress could be achieved over reasonable time scales. MACS amplifies a natural feedback in the Earth system for which comparable analogs exist, and its integration within existing agricultural and economic processes provides opportunities for global participation and co-benefits in a future MACS industry. Marine anoxic carbon storage (MACS) of lignocellulosic biomass is therefore worthy of serious consideration as an important potential CDR pathway for climate mitigation in coming decades.

Code, data, or code and data availability

No new data or code are used in this manuscript.

Author contributions

All authors participated in the in-person Bucharest Workshop discussions except TT, DS, NE, and SH. All authors contributed to the analysis and content of the manuscript. MR led writing and preparation of the manuscript.

Competing interests

In addition to her primary role as UCSB faculty, MR serves as the Chief Science Officer for Carboniferous, a U.S.-based startup company exploring potential applications of MACS. Both the coordination of the Bucharest Workshop and the resulting analysis presented here were supported by philanthropic funds and/or grants to UCSB; an unconflicted PI is included on all CDR-related funds in the Raven group.

Co-author TT is a member of the editorial team at Biogeosciences.

Co-authors TL and SH are employed at Puro.earth Oy, a Finland-based crediting platform for engineered carbon removals whose portfolio includes a standard for Marine Anoxic Carbon Storage.

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Voynova Y. G., Macovei V., Slabakova V., Drumeva N., Balan S., Neumann A. (in prep). Increasing temperatures and storm impact on the Black Sea oxygen deficit and carbon flux: the changing Black Sea, lessons from Storm Daniel.

Zarokanellos N. D., Ungureanu V. G., Papathanassiou V., Tyler A., Stanica A., and Tintoré J. (in prep). Revealing restratification processes and their biogeochemical impact in the Black Sea: Insights from Underwater Glider Observations.

References:

Addy S. K. and Behrens E. W. (1980) Time of accumulation of hypersaline anoxic brine in Orca basin (Gulf of Mexico). *Marine Geology* **37**, 241–252. doi:10.1016/0025-3227(80)90104-8

American Geophysical Union (2024) Ethical Framework Principles for Climate Intervention Research. doi:10.22541/essoar.172917365.53105072/v1

Armstrong McKay D. I., Staal A., Abrams J. F., Winkelmann R., Sakschewski B., Loriani S., Fetzer I., Cornell S. E., Rockström J. and Lenton T. M. (2022) Exceeding 1.5°C global warming could trigger multiple climate tipping points. *Science* **377**, eabn7950. doi:10.1126/science.abn7950

Arndt S., Jørgensen B. B., LaRowe D. E., Middelburg J. J., Pancost R. D. and Regnier P. (2013) Quantifying the degradation of organic matter in marine sediments: A review and synthesis. *Earth-Science Reviews* **123**, 53–86. doi:10.1016/j.earscirev.2013.02.008

Baker M. L., Hage S., Talling P. J., Acikalin S., Hilton R. G., Haghipour N., Ruffell S. C., Pope E. L., Jacinto R. S., Clare M. A. and Sahin S. (2024) Globally significant mass of terrestrial organic carbon efficiently transported by canyon-flushing turbidity currents. *Geology* **52**, 631–636. doi:10.1130/G51976.1

Ballantyne A. P., Alden C. B., Miller J. B., Tans P. P. and White J. W. C. (2012) Increase in observed net carbon dioxide uptake by land and oceans during the past 50 years. *Nature* **488**, 70–72. doi:10.1038/nature11299

Bar-On Y. M., Li X., O'Sullivan M., Wigneron J.-P., Sitch S., Ciais P., Frankenberg C. and Fischer W. W. (2025) Recent gains in global terrestrial carbon stocks are mostly stored in nonliving pools. *Science* **387**, 1291–1295. doi:10.1126/science.adk1637

Battin T. J., Lauerwald R., Bernhardt E. S., Bertuzzo E., Gener L. G., Hall R. O., Hotchkiss E. R., Maavara T., Pavelsky T. M., Ran L., Raymond P., Rosentreter J. A. and Regnier P. (2023) River ecosystem metabolism and carbon biogeochemistry in a changing world. *Nature* **613**, 449–459. doi:10.1038/s41586-022-05500-8

Berger M., Kwiatkowski L., Ho D. T. and Bopp L. (2023) Ocean dynamics and biological feedbacks limit the potential of macroalgae carbon dioxide removal. *Environ. Res. Lett.* **18**, 024039. doi:10.1088/1748-9326/acb06e

Blair N. E. and Aller R. C. (2012) The Fate of Terrestrial Organic Carbon in the Marine Environment. *Annu. Rev. Mar. Sci.* **4**, 401–423. doi:10.1146/annurev-marine-120709-142717

Brewer P. G. and Murray J. W. (1973) Carbon, nitrogen and phosphorus in the black sea. *Deep Sea Research and Oceanographic Abstracts* **20**, 803–818. doi:10.1016/0011-7471(73)90003-X

Buchanan P. J., Sun X., Weissman J., McCoy D., Bianchi D. and Zakem E. J. (2023) Oxygen intrusions sustain aerobic nitrite-oxidizing bacteria in anoxic marine zones. doi:10.1101/2023.02.22.529547

Burdige D. J. (2005) Burial of terrestrial organic matter in marine sediments: A re-assessment. *Global Biogeochemical Cycles* **19**, 2004GB002368. doi:10.1029/2004GB002368

Capet A., Beckers J.-M. and Grégoire M. (2013) Drivers, mechanisms and long-term variability of seasonal hypoxia on the Black Sea northwestern shelf – is there any recovery after eutrophication? *Biogeosciences* **10**, 3943–3962. doi:10.5194/bg-10-3943-2013

Capet A., Stanev E. V., Beckers J.-M., Murray J. W. and Grégoire M. (2016) Decline of the Black Sea oxygen inventory. *Biogeosciences* **13**, 1287–1297. doi:10.5194/bg-13-1287-2016

Cohen Y. and Gordon L. I. (1978) Nitrous oxide in the oxygen minimum of the eastern tropical North Pacific: evidence for its consumption during denitrification and possible mechanisms for its production. *Deep Sea Research* **25**, 509–524. doi:10.1016/0146-6291(78)90640-9

Contreras-Rosales L. A., Schefuß E., Meyer V., Palamenghi L., Lückge A. and Jennerjahn T. C. (2016) Origin and fate of sedimentary organic matter in the northern Bay of Bengal during the last 18 ka. *Global and Planetary Change* **146**, 53–66. doi:10.1016/j.gloplacha.2016.09.008

Cooley S. R., Klinsky S., Morrow D. R. and Satterfield T. (2023) Sociotechnical Considerations About Ocean Carbon Dioxide Removal. *Annu. Rev. Mar. Sci.* **15**, 41–66. doi:10.1146/annurev-marine-032122-113850

Curray J. R., Emmel F. J. and Moore D. G. (2002) The Bengal Fan: morphology, geometry, stratigraphy, history and processes. *Marine and Petroleum Geology* **19**, 1191–1223. doi:10.1016/S0264-8172(03)00035-7

Damien P., Sheinbaum J., Pasqueron De Fommervault O., Jouanno J., Linacre L. and Duteil O. (2021) Do Loop Current eddies stimulate productivity in the Gulf of Mexico? *Biogeosciences* **18**, 4281–4303. doi:10.5194/bg-18-4281-2021

Demidov A. B. (2008) Seasonal dynamics and estimation of the annual primary production of phytoplankton in the Black Sea. *Oceanology* **48**, 664–678. doi:10.1134/S0001437008050068

Deutsch C., Penn J. L. and Lucey N. (2024) Climate, Oxygen, and the Future of Marine Biodiversity. *Annu. Rev. Mar. Sci.* **16**, 217–245. doi:10.1146/annurev-marine-040323-095231

Doney S. C. (2010) The Growing Human Footprint on Coastal and Open-Ocean Biogeochemistry. *Science* **328**, 1512–1516. doi:10.1126/science.1185198

Drijfhout S., Angevaare J. R., Mecking J., Van Westen R. M. and Rahmstorf S. (2025) Shutdown of northern Atlantic overturning after 2100 following deep mixing collapse in CMIP6 projections. *Environ. Res. Lett.* **20**, 094062. doi:10.1088/1748-9326/adfa3b

Dunne J. P., Sarmiento J. L. and Gnanadesikan A. (2007) A synthesis of global particle export from the surface ocean and cycling through the ocean interior and on the seafloor: Global particle export and cycling. *Global Biogeochem. Cycles* **21**, 1–16. doi:10.1029/2006GB002907

Dyrssen D. and Kremling K. (1990) Increasing hydrogen sulfide concentration and trace metal behavior in the anoxic Baltic waters. *Marine Chemistry* **30**, 193–204. doi:10.1016/0304-4203(90)90070-S

France-Lanord C. and Derry L. A. (1997) Organic carbon burial forcing of the carbon cycle from Himalayan erosion. *Nature* **390**, 65–67. doi:10.1038/36324

Friedlingstein P., O'Sullivan M., Jones M. W., Andrew R. M., Hauck J., Landschützer P., Le Quéré C., Li H., Luijkx I. T., Olsen A., Peters G. P., Peters W., Pongratz J., Schwingshackl C., Sitch S., Canadell J. G., Ciais P., Jackson R. B., Alin S. R., Arneth A., Arora V., Bates N. R., Becker M., Bellouin N., Berghoff C. F., Bittig H. C., Bopp L., Cadule

P., Campbell K., Chamberlain M. A., Chandra N., Chevallier F., Chini L. P., Colligan T., Decayeux J., Djeutchouang L. M., Dou X., Duran Rojas C., Enyo K., Evans W., Fay A. R., Feely R. A., Ford D. J., Foster A., Gasser T., Gehlen M., Gkritzalis T., Grassi G., Gregor L., Gruber N., Gürses Ö., Harris I., Hefner M., Heinke J., Hurtt G. C., Iida Y., Ilyina T., Jacobson A. R., Jain A. K., Jarníková T., Jersild A., Jiang F., Jin Z., Kato E., Keeling R. F., Klein Goldewijk K., Knauer J., Korsbakken J. I., Lan X., Lauvset S. K., Lefèvre N., Liu Z., Liu J., Ma L., Maksyutov S., Marland G., Mayot N., McGuire P. C., Metzl N., Monacci N. M., Morgan E. J., Nakaoka S.-I., Neill C., Niwa Y., Nützel T., Olivier L., Ono T., Palmer P. I., Pierrot D., Qin Z., Resplandy L., Roobaert A., Rosan T. M., Rödenbeck C., Schwinger J., Smallman T. L., Smith S. M., Sospedra-Alfonso R., Steinhoff T., Sun Q., Sutton A. J., Séférian R., Takao S., Tatebe H., Tian H., Tilbrook B., Torres O., Tourigny E., Tsujino H., Tubiello F., Van Der Werf G., Wanninkhof R., Wang X., Yang D., Yang X., Yu Z., Yuan W., Yue X., Zaehle S., Zeng N. and Zeng J. (2025) Global Carbon Budget 2024. *Earth Syst. Sci. Data* **17**, 965–1039. doi:10.5194/essd-17-965-2025

Galy V., France-Lanord C., Beyssac O., Faure P., Kudrass H. and Palhol F. (2007) Efficient organic carbon burial in the Bengal fan sustained by the Himalayan erosional system. *Nature* **450**, 407–410. doi:10.1038/nature06273

Ganti G., Gasser T., Bui M., Geden O., Lamb W. F., Minx J. C., Schleussner C.-F. and Gidden M. J. (2024) Evaluating the near- and long-term role of carbon dioxide removal in meeting global climate objectives. *Commun Earth Environ* **5**, 377. doi:10.1038/s43247-024-01527-z

GESAMP (2025) *The State of the Science for Marine Carbon Dioxide Removal (mCDR) – A Scientific Summary for Policy-Makers*., UNESCO-IOC, Paris. doi:10.5281/ZENODO.15490407

Gilly W. F., Beman J. M., Litvin S. Y. and Robison B. H. (2013) Oceanographic and Biological Effects of Shoaling of the Oxygen Minimum Zone. *Annu. Rev. Mar. Sci.* **5**, 393–420. doi:10.1146/annurev-marine-120710-100849

Grilli R., Birot D., Schumacher M., Paris J.-D., Blouzon C., Donval J. P., Guyader V., Leau H., Giunta T., Delmotte M., Radulescu V., Balan S., Greinert J. and Ruffine L. (2021) Inter-Comparison of the Spatial Distribution of Methane in the Water Column From Seafloor Emissions at Two Sites in the Western Black Sea Using a Multi-Technique Approach. *Front. Earth Sci.* **9**, 626372. doi:10.3389/feart.2021.626372

Haas S. (2020) Biogeochemical transformations under contrasting mixing regimes in two coastal water columns. PhD thesis, Dalhousie University, Department of Oceanography.

Haeckel M., Bialas J., Klaucke I., Wallmann K., Bohrmann G. and Schwalenberg K. (2015) Gas hydrate occurrences in the Black sea – new observations from the German SUGAR Project. *Fire in the Ice: Methane Hydrate Newsletter*, 6–9.

Hage S., Baker M. L., Babonneau N., Soulet G., Dennielou B., Silva Jacinto R., Hilton R. G., Galy V., Baudin F., Rabouille C., Vic C., Sahin S., Açikalin S. and Talling P. J. (2024) How is particulate organic carbon transported through the river-fed submarine Congo Canyon to the deep sea? *Biogeosciences* **21**, 4251–4272. doi:10.5194/bg-21-4251-2024

Hage S., Galy V. V., Cartigny M. J. B., Acikalin S., Clare M. A., Gröcke D. R., Hilton R. G., Hunt J. E., Lintern D. G., McGhee C. A., Parsons D. R., Stacey C. D., Sumner E. J. and Talling P. J. (2020) Efficient preservation of young terrestrial organic carbon in sandy turbidity-current deposits. *Geology* **48**, 882–887. doi:10.1130/G47320.1

Harvey H. R. and Kennicutt M. C. (1992) Selective alteration of Sargassum lipids in anoxic sediments of the Orca Basin. *Organic Geochemistry* **18**, 181–187. doi:10.1016/0146-6380(92)90128-K

Hebbeln D., Wienberg C., Dullo W.-C., Freiwald A., Mienis F., Orejas C. and Titschack J. (2020) Cold-water coral reefs thriving under hypoxia. *Coral Reefs* **39**, 853–859. doi:10.1007/s00338-020-01934-6

https://doi.org/10.5194/egusphere-2025-6086
Preprint. Discussion started: 18 December 2025

Hinrichs K.-U. and Boetius A. (2002) The Anaerobic Oxidation of Methane: New Insights in Microbial Ecology and Biogeochemistry. In *Ocean Margin Systems* (eds. G. Wefer, D. Billett, D. Hebbeln, B. B. Jørgensen, M. Schlüter, and T. C. E. Van Weering). Springer Berlin Heidelberg, Berlin, Heidelberg. pp. 457–477. doi:10.1007/978-3-662-05127-6_28

Hiscock W. T. and Millero F. J. (2006) Alkalinity of the anoxic waters in the Western Black Sea. *Deep Sea Research Part II: Topical Studies in Oceanography* **53**, 1787–1801. doi:10.1016/j.dsr2.2006.05.020

Huang Y., Freeman K. H., Wilkin R. T., Arthur M. A. and Jones A. D. (2000) Black Sea chemocline oscillations during the Holocene: molecular and isotopic studies of marginal sediments. *Organic Geochemistry* **31**, 1525–1531. doi:10.1016/S0146-6380(00)00090-5

Huntington T., Scown C., Breunig H., Kavvada O., Cui X., Hendrickson T., Nordahl S., Baral N. R., Yang M. and Moore M. (2024) BioSiting Tool (BioSiting) v2. doi:10.11578/DC.20240815.9

Hurtgen M. T., Lyons T., Ingall E. and Cruse A. (1999) Anomalous enrichments of iron monosulfide in euxinic marine sediments and the role of H 2 S in iron sulfide transformations; examples from Effingham Inlet, Orca Basin, and the Black Sea. *American Journal of Science* **299**, 556–588. doi:10.2475/ajs.299.7-9.556

Inglis G., Hemingway J., Stockey R., Hollingsworth E., Valdes P., Farnsworth A., Elling F., Badger M., Pancost R., Pearson A., Gernon T. and Hilton R. (2025) Marine burial of terrestrial organic carbon modulates past warm climates. doi:10.21203/rs.3.rs-7272686/v1

IPCC AR6 (2021) Intergovernmental Panel on Climate Change, Working Group II, Sixth Assessment Report: Technical Summary.

Jarvis I., Lignum J. S., Gröcke D. R., Jenkyns H. C. and Pearce M. A. (2011) Black shale deposition, atmospheric CO2 drawdown, and cooling during the Cenomanian-Turonian Oceanic Anoxic Event: CO2 drawdown and cooling during OAE2. *Paleoceanography* **26**, n/a-n/a. doi:10.1029/2010PA002081

Keil R. G., Nuwer J. M. and Strand S. E. (2010) Burial of agricultural byproducts in the deep sea as a form of carbon sequestration: A preliminary experiment. *Marine Chemistry* **122**, 91–95. doi:10.1016/j.marchem.2010.07.007

Kennett J. P. and Penrose N. L. (1978) Fossil Holocene seaweed and attached calcareous polychaetes in an anoxic basin, Gulf of Mexico. *Nature* **276**, 172–173. doi:10.1038/276172a0

Kessler J. D., Reeburgh W. S., Southon J., Seifert R., Michaelis W. and Tyler S. C. (2006) Basin-wide estimates of the input of methane from seeps and clathrates to the Black Sea. *Earth and Planetary Science Letters* **243**, 366–375. doi:10.1016/j.epsl.2006.01.006

Kessler J. D., Valentine D. L., Redmond M. C., Du M., Chan E. W., Mendes S. D., Quiroz E. W., Villanueva C. J., Shusta S. S., Werra L. M., Yvon-Lewis S. A. and Weber T. C. (2011) A Persistent Oxygen Anomaly Reveals the Fate of Spilled Methane in the Deep Gulf of Mexico. *Science* **331**, 312–315. doi:10.1126/science.1199697

Kim K.-R. and Craig H. (1990) Two-isotope characterization of N20 in the Pacific Ocean and constraints on its origin in deep water. *Nature* **347**, 58–61. doi:10.1038/347058a0

Köhn E. E., Münnich M., Vogt M., Desmet F. and Gruber N. (2022) Strong Habitat Compression by Extreme Shoaling Events of Hypoxic Waters in the Eastern Pacific. *JGR Oceans* **127**, e2022JC018429. doi:10.1029/2022JC018429

Konovalov S. K. and Murray J. W. (2001) Variations in the chemistry of the Black Sea on a time scale of decades (1960–1995). *Journal of Marine Systems* **31**, 217–243.

Konovalov S. K., Murray J. W., Luther G. W. and Tebo B. M. (2006) Processes controlling the redox budget for the oxic/anoxic water column of the Black Sea. *Deep Sea Research Part II: Topical Studies in Oceanography* **53**, 1817–1841. doi:10.1016/j.dsr2.2006.03.013

Kroiss H., Zessner M. and Lampert C. (2006) daNUbs: Lessons learned for nutrient management in the Danube Basin and its relation to Black Sea euthrophication. *Chemistry and Ecology* **22**, 347–357. doi:10.1080/02757540600917518

Lalonde K., Mucci A., Ouellet A. and Gélinas Y. (2012) Preservation of organic matter in sediments promoted by iron. *Nature* **483**, 198–200. doi:10.1038/nature10855

Langholtz M. (2024) 2023 Billion-Ton Report. doi:10.23720/BT2023/2316165

Lee B.-S., Bullister J. L., Murray J. W. and Sonnerup R. E. (2002) Anthropogenic chlorofluorocarbons in the Black Sea and the Sea of Marmara. *Deep Sea Research Part I: Oceanographic Research Papers* **49**, 895–913. doi:10.1016/S0967-0637(02)00005-5

Ludwig W., Dumont E., Meybeck M. and Heussner S. (2009) River discharges of water and nutrients to the Mediterranean and Black Sea: Major drivers for ecosystem changes during past and future decades? *Progress in Oceanography* **80**, 199–217. doi:10.1016/j.pocean.2009.02.001

Masson D. G., Huvenne V. A. I., De Stigter H. C., Wolff G. A., Kiriakoulakis K., Arzola R. G. and Blackbird S. (2010) Efficient burial of carbon in a submarine canyon. *Geology* **38**, 831–834. doi:10.1130/G30895.1

McGinnis D. F., Greinert J., Artemov Y., Beaubien S. E. and Wüest A. (2006) Fate of rising methane bubbles in stratified waters: How much methane reaches the atmosphere? *J. Geophys. Res.* **111**, 2005JC003183. doi:10.1029/2005JC003183

Middelburg J. J., Soetaert K. and Hagens M. (2020) Ocean Alkalinity, Buffering and Biogeochemical Processes. *Rev. Geophys.* **58**. doi:10.1029/2019RG000681

Mikaelyan A. S., Zatsepin A. G. and Chasovnikov V. K. (2013) Long-term changes in nutrient supply of phytoplankton growth in the Black Sea. *Journal of Marine Systems* **117–118**, 53–64. doi:10.1016/j.jmarsys.2013.02.012

Müller B., Meyer J. S. and Gächter R. (2016) Alkalinity regulation in calcium carbonate-buffered lakes: Alkalinity regulation in calcium carbonate-buffered lakes. *Limnol. Oceanogr.* **61**, 341–352. doi:10.1002/lno.10213

Müller C., Cramer W., Hare W. L. and Lotze-Campen H. (2011) Climate change risks for African agriculture. *Proc. Natl. Acad. Sci. U.S.A.* **108**, 4313–4315. doi:10.1073/pnas.1015078108

Murray J. W., Top Z. and Özsoy E. (1991) Hydrographic properties and ventilation of the Black Sea. *Deep Sea Research Part A. Oceanographic Research Papers* **38**, S663–S689. doi:10.1016/S0198-0149(10)80003-2

Nabeh N., Brokaw C. and Picard A. (2022) Quantification of Organic Carbon Sequestered by Biogenic Iron Sulfide Minerals in Long-Term Anoxic Laboratory Incubations. *Front. Microbiol.* **13**, 662219. doi:10.3389/fmicb.2022.662219

National Academies of Sciences, Engineering, and Medicine, Committee on A Research Strategy for Ocean-based Carbon Dioxide Removal and Sequestration, Ocean Studies Board, and Division on Earth and Life Studies (2022) *A Research*

Strategy for Ocean-based Carbon Dioxide Removal and Sequestration., National Academies Press, Washington, D.C. doi:10.17226/26278

Nawaz S., Peterson St-Laurent G. and Satterfield T. (2023) Public evaluations of four approaches to ocean-based carbon dioxide removal. *Climate Policy* **23**, 379–394. doi:10.1080/14693062.2023.2179589

Nigro L. M., Elling F. J., Hinrichs K.-U., Joye S. B. and Teske A. (2020) Microbial ecology and biogeochemistry of hypersaline sediments in Orca Basin ed. W. J. Brazelton. *PLoS ONE* **15**, e0231676. doi:10.1371/journal.pone.0231676

Oguz T. and Gilbert D. (2007) Abrupt transitions of the top-down controlled Black Sea pelagic ecosystem during 1960–2000: Evidence for regime-shifts under strong fishery exploitation and nutrient enrichment modulated by climate-induced variations. *Deep Sea Research Part I: Oceanographic Research Papers* **54**, 220–242. doi:10.1016/j.dsr.2006.09.010

Owens J. D., Lyons T. W. and Lowery C. M. (2018) Quantifying the missing sink for global organic carbon burial during a Cretaceous oceanic anoxic event. *Earth and Planetary Science Letters* **499**, 83–94. doi:10.1016/j.epsl.2018.07.021

Oyedeji O. A., Moore A. M., Kamath D., Busch I. K., Curran S. J., Henriksen M., Jamieson M., Fout T. E., Webb E. G., Langholtz M. H. and Theiss T. J. (2025) Economic and environmental performance of biomass gasification for renewable natural gas production in the context of the U.S. natural gas supply. *Biomass and Bioenergy* **197**, 107670. doi:10.1016/j.biombioe.2025.107670

Pacheco-Ruiz R., Adams J., Pedrotti F., Grant M., Holmlund J. and Bailey C. (2019) Deep sea archaeological survey in the Black Sea – Robotic documentation of 2,500 years of human seafaring. *Deep Sea Research Part I: Oceanographic Research Papers* **152**, 103087. doi:10.1016/j.dsr.2019.103087

Pakhomova S., Vinogradova E., Yakushev E., Zatsepin A., Shtereva G., Chasovnikov V. and Podymov O. (2014) Interannual variability of the Black Sea Proper oxygen and nutrients regime: The role of climatic and anthropogenic forcing. *Estuarine, Coastal and Shelf Science* **140**, 134–145. doi:10.1016/j.ecss.2013.10.006

Papadomanolaki N. M., Lenstra W. K., Wolthers M. and Slomp C. P. (2022) Enhanced phosphorus recycling during past oceanic anoxia amplified by low rates of apatite authigenesis. *Sci. Adv.* **8**, eabn2370. doi:10.1126/sciadv.abn2370

Pett-Ridge J., Kuebbing S., Mayer A., Hovorka S., Pilorgé H., Baker S., Pang S., Scown C., Mayfield K., Wong A., Aines R., Ammar H., Aui A., Ashton M., Basso B., Bradford M., Bump A., Busch I., Calzado E., Chirigotis J., Clauser N., Crotty S., Dahl N., Dai T., Ducey M., Dumortier J., Ellebracht N., Egui R., Fowler A., Georgiou K., Giannopoulos D., Goldstein H., Harris T., Hayes D., Hellwinckel C., Ho A., Hong M., Hunter-Sellars E., Kirkendall W., Langholtz M., Layer M., Lee I., Lewis R., Li W., Liu W., Lozano J., Lunstrum A., McNeil W., Nico P., O'Rourke A., Paustian K., Peridas G., Pisciotta M., Price L., Psarras P., Robertson G., Sagues W., Sanchez D., Schmidt B., Slessarev E., Sokol N., Stanley A., Swan A., Toureene C., Wright M., Yao Y., Zhang B. and Zhang Y. (2023) *Roads to Removal: Options for Carbon Dioxide Removal in the United States*., doi:10.2172/2301853

Puro.earth (2025) Biomass Sourcing Criteria v1.2.

Puro.Earth (2025) *Terrestrial Storage of Biomass: Methodology for CO2 Removal*.,

Rao C. K., Naqvi S. W. A., Kumar M. D., Varaprasad S. J. D., Jayakumar D. A., George M. D. and Singbal S. Y. S. (1994) Hydrochemistry of the Bay of Bengal: possible reasons for a different water-column cycling of carbon and nitrogen from the Arabian Sea. *Marine Chemistry* **47**, 279–290. doi:10.1016/0304-4203(94)90026-4

Raven M. R., Crotteau M. A., Evans N., Girard Z. C., Martinez A. M., Young I. and Valentine D. L. (2024) Biomass Storage in Anoxic Marine Basins: Initial Estimates of Geochemical Impacts and CO_2 Sequestration Capacity. *AGU Advances* **5**, e2023AV000950. doi:10.1029/2023AV000950

Raven M. R., Evans N., Martinez A. M. and Phillips A. A. (2025) Big decisions from small experiments: observational strategies for biomass-based marine carbon storage. *Environ. Res. Lett.* **20**, 051001. doi:10.1088/1748-9326/adc28d

Raymond P. A., Saiers J. E. and Sobczak W. V. (2016) Hydrological and biogeochemical controls on watershed dissolved organic matter transport: pulse-shunt concept. *Ecology* **97**, 5–16. doi:10.1890/14-1684.1

Reeburgh W. S., Ward B. B., Whalen S. C., Sandbeck K. A., Kilpatrickt K. A. and Kerkhof L. J. (1991) Black Sea methane geochemistry. *Deep Sea Research Part A. Oceanographic Research Papers* **38**, S1189–S1210. doi:10.1016/S0198-0149(10)80030-5

Roberts K. E., Rohr T. W., Raven M. R., Diamond M. S., Visioni D., Heneghan R., Bianchi D., Ortega-Cisneros K., Morrison M. A., Heerden V. V., Wiseman N., Anil G., Cannizzo Z. J., Coll M., Coupe J., Freedman R. M., Kravitz B., Krumhardt K. M., Kwiatkowski L., Lovenduski N. S., Luo J. Y., Olivarez H. C., Petrik C. M., Robock A., Steenbeek J. G. and Harrison C. S. (2025) Potential impacts of climate interventions on marine ecosystems. doi:10.22541/essoar.173482381.11257447/v1

Rosa L., Sanchez D. L. and Mazzotti M. (2021) Assessment of carbon dioxide removal potential *via* BECCS in a carbon-neutral Europe. *Energy Environ. Sci.* **14**, 3086–3097. doi:10.1039/D1EE00642H

Saager P. M., Schijf J. and De Baar H. J. W. (1993) Trace-metal distributions in seawater and anoxic brines in the eastern Mediterranean Sea. *Geochimica et Cosmochimica Acta* **57**, 1419–1432. doi:10.1016/0016-7037(93)90003-F

Sackett W. M., Brooks J. M., Bernard B. B., Schwab C. R., Chung H. and Parker R. A. (1979) A carbon inventory for Orca Basin brines and sediments. *Earth and Planetary Science Letters* **44**, 73–81. doi:10.1016/0012-821X(79)90009-8

Sanchez D. L., Psarras P., Murnen H. K. and Rogers B. (2025) Carbon removal efficiency and energy requirement of engineered carbon removal technologies. *RSC Sustain.*, 10.1039.D4SU00552J. doi:10.1039/D4SU00552J

Schmale O., Haeckel M. and McGinnis D. F. (2011) Response of the Black Sea methane budget to massive short-term submarine inputs of methane. *Biogeosciences* **8**, 911–918. doi:10.5194/bg-8-911-2011

Schunck H., Lavik G., Desai D. K., Großkopf T., Kalvelage T., Löscher C. R., Paulmier A., Contreras S., Siegel H., Holtappels M., Rosenstiel P., Schilhabel M. B., Graco M., Schmitz R. A., Kuypers M. M. M. and LaRoche J. (2013) Giant Hydrogen Sulfide Plume in the Oxygen Minimum Zone off Peru Supports Chemolithoautotrophy ed. L. J. Stal. *PLoS ONE* **8**, e68661. doi:10.1371/journal.pone.0068661

Sergeev D. A., Troitskaya Y. I., Ermakova O. S. and Orekhova N. A. (2024) Investigation of the Relationship between Partial Pressure of Carbon Dioxide and Sea Surface Temperature in the Cyclic Seasonal Variations in the Black Sea. *Morskoy Gidrofizicheskiy Zhurnal* **40**, 805–820. doi:10.2172/10107773

Shah S. R., Joye S. B., Brandes J. A. and McNichol A. P. (2013) Carbon isotopic evidence for microbial control of carbon supply to Orca Basin at the seawater–brine interface. *Biogeosciences* **10**, 3175–3183. doi:10.5194/bg-10-3175-2013

Shokes R. F., Trabant P. K., Presley B. J. and Reid D. F. (1977) Anoxic, Hypersaline Basin in the Northern Gulf of Mexico. *Science* **196**, 1443–1446. doi:10.1126/science.196.4297.1443

Silverman-Roati K. and Webb R. M. (2025) International Legal Guidelines for Marine Carbon Dioxide Removal Governance under the London Convention and London Protocol.

Snyder B. F. (2022) Biomass slurry fracture injection as a potential low-cost negative emissions technology. *Environ. Res. Lett.* **17**, 024013. doi:10.1088/1748-9326/ac4c5c

Stanev E. V., Wahle K. and Staneva J. (2022) The Synergy of Data From Profiling Floats, Machine Learning and Numerical Modeling: Case of the Black Sea Euphotic Zone. *JGR Oceans* **127**. doi:10.1029/2021JC018012

Starostenko V. I., Rusakov O. M., Shnyukov E. F., Kobolev V. P. and Kutas R. I. (2010) Methane in the northern Black Sea: characterization of its geomorphological and geological environments. *SP* **340**, 57–75. doi:10.1144/SP340.5

Strand S. E. and Benford G. (2009) Ocean Sequestration of Crop Residue Carbon: Recycling Fossil Fuel Carbon Back to Deep Sediments. *Environ. Sci. Technol.* **43**, 1000–1007. doi:10.1021/es8015556

Tagliabue A., Twining B. S., Barrier N., Maury O., Berger M. and Bopp L. (2023) Ocean iron fertilization may amplify climate change pressures on marine animal biomass for limited climate benefit. *Global Change Biology*, gcb.16854. doi:10.1111/gcb.16854

Talling P. J., Baker M. L., Pope E. L., Ruffell S. C., Jacinto R. S., Heijnen M. S., Hage S., Simmons S. M., Hasenhündl M., Heerema C. J., McGhee C., Apprioual R., Ferrant A., Cartigny M. J. B., Parsons D. R., Clare M. A., Tshimanga R. M., Trigg M. A., Cula C. A., Faria R., Gaillot A., Bola G., Wallance D., Griffiths A., Nunny R., Urlaub M., Peirce C., Burnett R., Neasham J. and Hilton R. J. (2022) Longest sediment flows yet measured show how major rivers connect efficiently to deep sea. *Nat Commun* **13**, 4193. doi:10.1038/s41467-022-31689-3

Talling P. J., Hage S., Baker M. L., Bianchi T. S., Hilton R. G. and Maier K. L. (2024) The Global Turbidity Current Pump and Its Implications for Organic Carbon Cycling. *Annu. Rev. Mar. Sci.* **16**, 105–133. doi:10.1146/annurev-marine-032223-103626

Tiano L., Garcia-Robledo E., Dalsgaard T., Devol A. H., Ward B. B., Ulloa O., Canfield D. E. and Peter Revsbech N. (2014) Oxygen distribution and aerobic respiration in the north and south eastern tropical Pacific oxygen minimum zones. *Deep Sea Research Part I: Oceanographic Research Papers* **94**, 173–183. doi:10.1016/j.dsr.2014.10.001

Van Cappellen P., Viollier E., Roychoudhury A., Clark L., Ingall E., Lowe K. and Dichristina T. (1998) Biogeochemical Cycles of Manganese and Iron at the Oxic−Anoxic Transition of a Stratified Marine Basin (Orca Basin, Gulf of Mexico). *Environ. Sci. Technol.* **32**, 2931–2939. doi:10.1021/es980307m

Vaquer-Sunyer R. and Duarte C. M. (2008) Thresholds of hypoxia for marine biodiversity. *Proc. Natl. Acad. Sci. U.S.A.* **105**, 15452–15457. doi:10.1073/pnas.0803833105

Vaughan N., Fuss S., Buck H., Schenuit F., Pongratz J., Schulte I., Lamb W. F., Probst B., Edwards M., Nemet G. F., Cox E., Smith S., Injy Johnstone, Geden O., Burke J., Gidden M., Roe S., Müller-Hansen F. and Minx J. (2024) The State of Carbon Dioxide Removal - 2nd Edition. doi:10.17605/OSF.IO/F85QJ

Wanninkhof R., Lewis E., Feely R. A. and Millero F. J. (1999) The optimal carbonate dissociation constants for determining surface water pCO2 from alkalinity and total inorganic carbon. *Marine Chemistry* **65**, 291–301. doi:10.1016/S0304-4203(99)00021-3

Wenk C. B., Frame C. H., Koba K., Casciotti K. L., Veronesi M., Niemann H., Schubert C. J., Yoshida N., Toyoda S., Makabe A., Zopfi J. and Lehmann M. F. (2016) Differential N_2 O dynamics in two oxygen-deficient lake basins revealed by

1020 stable isotope and isotopomer distributions: Differential N_2O Dynamics in Two Lake Basins. *Limnol. Oceanogr.* **61**, 1735–1749. doi:10.1002/lno.10329

Westley M. B., Yamagishi H., Popp B. N. and Yoshida N. (2006) Nitrous oxide cycling in the Black Sea inferred from stable isotope and isotopomer distributions. *Deep Sea Research Part II: Topical Studies in Oceanography* **53**, 1802–1816. doi:10.1016/j.dsr2.2006.03.012

1025 Westrich J. T. and Berner R. A. (1984) The Role of Sedimentary Organic Matter in Bacterial Sulfate Reduction: The G Model Tested. *Limnology and Oceanography* **29**, 236–249.

Wiesenburg D. A., Brooks J. M. and Bernard B. B. (1985) Biogenic hydrocarbon gases and sulfate reduction in the Orca Basin brine. *Geochimica et Cosmochimica Acta* **49**, 2069–2080. doi:10.1016/0016-7037(85)90064-X

Zeng N., Zhao X., Poisson G., Clifford B., Liu Y., Liu H., Meng T., Picard L., Zeng-Mariotti E., Zaitchik B. and Hu L. (2024)
1030 3775-year-old wood burial supports "wood vaulting" as a durable carbon removal method. *Science* **385**, 1454–1459. doi:10.1126/science.adm8133

Zhang J.-Z. and Millero F. J. (1993) The chemistry of the anoxic waters in the Cariaco Trench. *Deep Sea Research Part I: Oceanographic Research Papers* **40**, 1023–1041. doi:10.1016/0967-0637(93)90088-K

Zhang Y., Li W., Sun G., Mao J., Dannenberg M., Xiao J., Li Z., Zhao H., Zhang Q., Hu S., Song C. and Cassar N. (2025)
1035 Contrasting biological production trends over land and ocean. *Nat. Clim. Chang.* **15**, 880–888. doi:10.1038/s41558-025-02375-1

Zhuang G.-C., Elling F. J., Nigro L. M., Samarkin V., Joye S. B., Teske A. and Hinrichs K.-U. (2016) Multiple evidence for methylotrophic methanogenesis as the dominant methanogenic pathway in hypersaline sediments from the Orca Basin, Gulf of Mexico. *Geochimica et Cosmochimica Acta* **187**, 1–20. doi:10.1016/j.gca.2016.05.005