

Supplement of

Secondary ice production affects tropical convective clouds under different aerosol conditions

5 **Mengyu Sun et al.**

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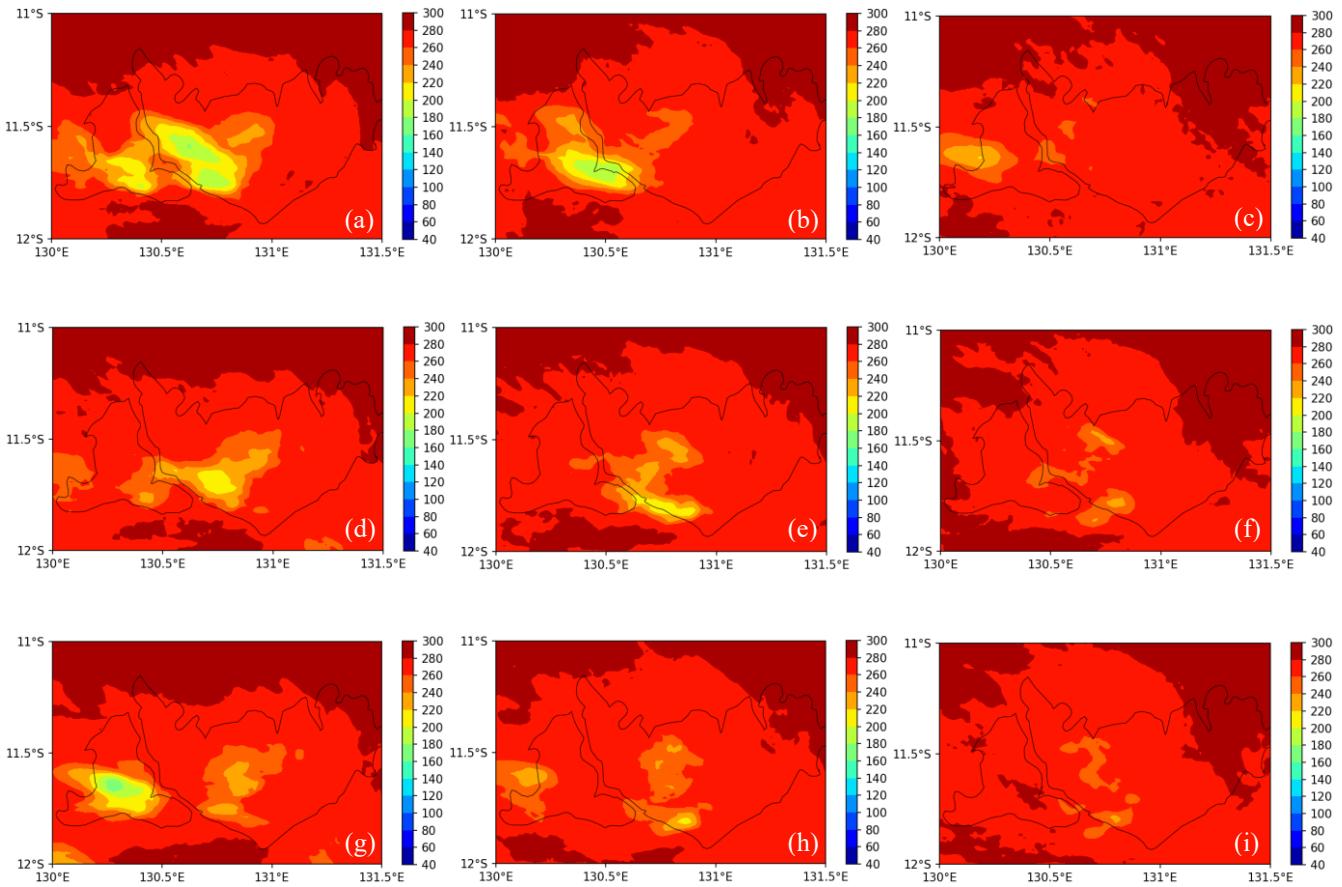


Figure S1. Spatial distribution of OLR (W m^{-2}) from noSIP simulations for (a–c) $N_d = 200 \text{ cm}^{-3}$, (d–f) $N_d = 400 \text{ cm}^{-3}$, and (g–i) $N_d = 800 \text{ cm}^{-3}$ at 06:30, 07:30, and 08:30 UTC on 6 February 2006.

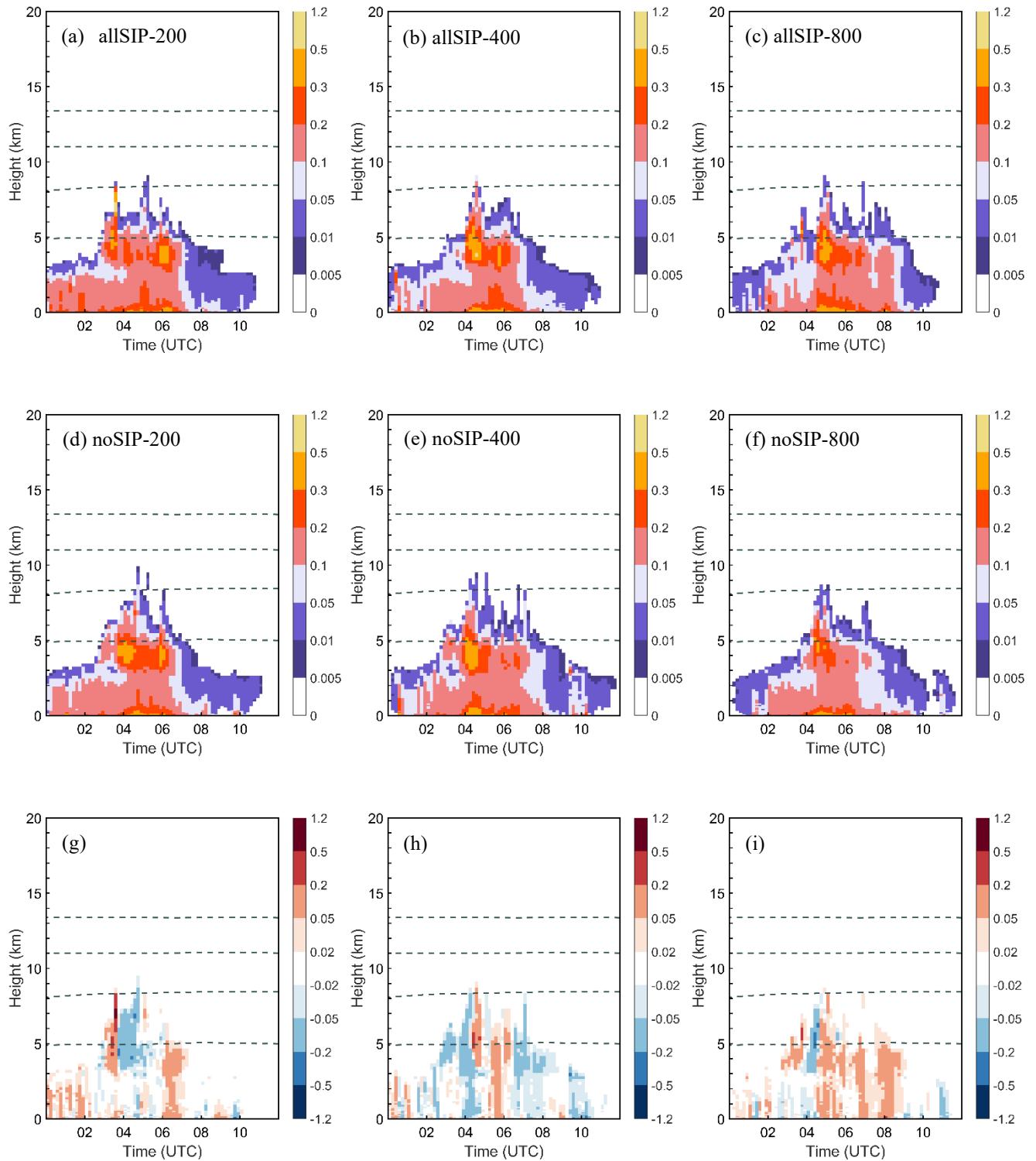


Figure S2. Time–height cross sections of rain water content (RWC; g m^{-3}) for simulations with all SIP mechanisms included: (a) allSIP-200, (b) allSIP-400, and (c) allSIP-800; and without SIP: (d) noSIP-200, (e) noSIP-400, and (f) noSIP-800. Panels (g–i) show the differences between allSIP and noSIP simulations at each CCN concentration ($N_d = 200, 400, 800 \text{ cm}^{-3}$; i.e., allSIP minus noSIP). Panels (a–f) are averaged over the area where either the ice water path or rain water path exceeds 1 g m^{-2} . The 0, -20, -40, and -60 °C isotherms are shown by the dashed lines.

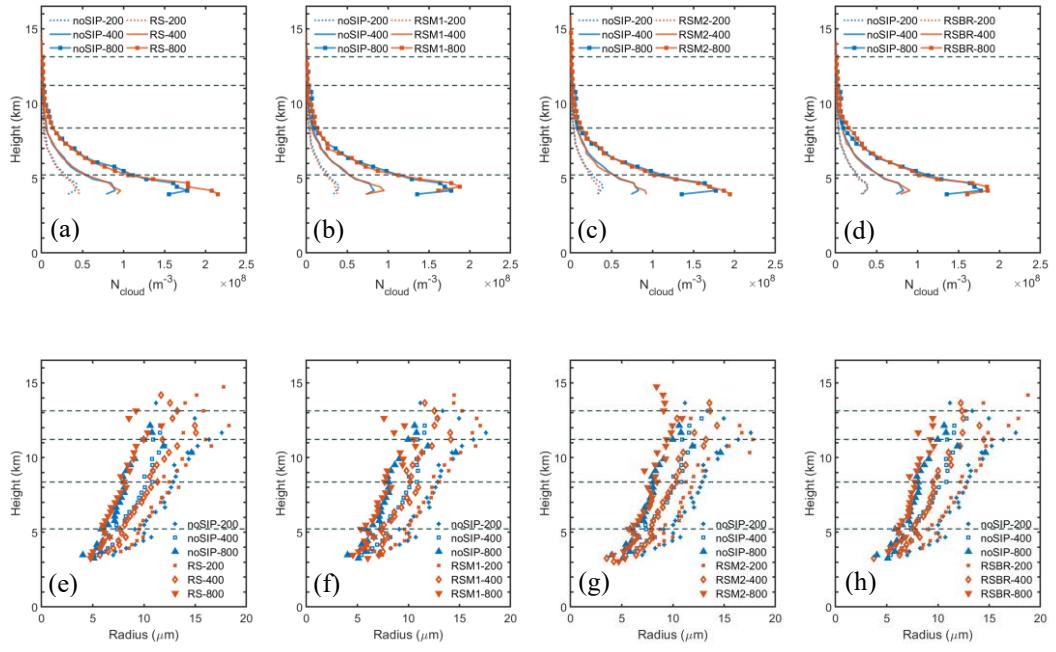


Figure S3. Vertical profiles of horizontally averaged (a-d) number concentration of cloud droplet (N_{cloud}) and (e-h) mean-mass radius of cloud droplet, from simulations without SIP (noSIP) and with different SIP configurations: (a, e) RS, (b, f) RSM1, (c, g) RSM2, and (d, h) RSBR. Dotted lines denote $N_d = 200 \text{ cm}^{-3}$, solid lines denote $N_d = 400 \text{ cm}^{-3}$, and solid lines with square markers denote $N_d = 800 \text{ cm}^{-3}$. All profiles are averaged over the area where either the ice water path or rain water path exceeds 1 g m^{-2} , during 02:30–08:30 UTC on 6 February 2006. Gray dashed lines indicate the 0, -20 , -40 , and -60 $^{\circ}\text{C}$ isotherms.

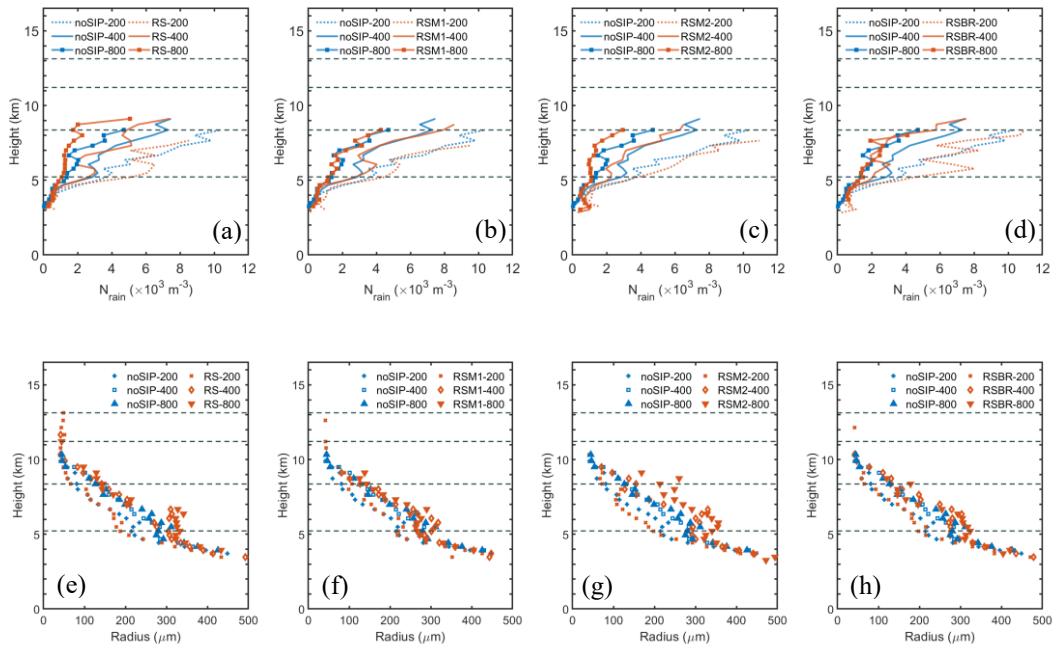


Figure S4. Same as Figure S3, but for (a-d) rain number concentration (N_{rain}), and (e-h) rain mean-mass radius.