

Monitoring the displacement of large alpine rock slope instabilities with L-band SAR interferometric techniques

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Abstract. Large rock slope instabilities develop over long periods and creep slowly over hundreds or thousands of years, until they undergo a 'slow to fast' evolution towards catastrophic collapse. Capturing this transition is key to manage related risks, especially considering ongoing climate change scenarios and human activities, that are expected to strongly influence geohazards. However, this is a challenging task due to the complexity of the underlying processes. Long-term, area-wide monitoring of slope movements is essential to understand landslide dynamics and evolution. Despite being widely used for landslide investigations, C-band SAR interferometry datasets suffer from decorrelation in vegetated areas and fast movements, limiting displacement retrieval in alpine regions. Emerging L-band systems, with reduced temporal decorrelation, can complement higher-frequency data by enabling measurements also in vegetated areas and capturing larger displacements. This work aims at analysing the potential benefits and limitations of L-band SAR interferometry applied to alpine landslide monitoring and at understanding if these data can help in mitigating current shortcomings of C-band SAR interferometry. We explore three different scenarios of large alpine slope instabilities in the European Alps, that threaten important economic and societal assets. We perform site-specific analysis, validation and interpretation of L-band SAR interferometry products derived from ALOS-2 PALSAR-2 and SAOCOM-1 satellite imagery, as well as of terrestrial data acquired by the GAMMA L-band SAR (GLSAR) instrument. Our results highlight the contributions of L-band InSAR products to the practical characterisation and interpretation of large rock slope instabilities and provide important recommendations for the recently launched L-band satellite SAR missions ALOS-4 PALSAR-3 and NISAR, as well as for the future L-band satellite SAR mission ROSE-L.

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30 **1 Introduction**

Large-scale rock slope instability phenomena are widespread in mountain environments. They significantly influence the long-term topographic, hydrological and sediment transfer characteristics of alpine slopes and have a high potential to threaten human lives, urban settlements, transport infrastructure and other lifelines (Crosta et al., 2013; Agliardi and Crippa, 2022). Ongoing changes in climatic conditions as well as human activity are expected to have a strong influence on geomorphological processes and to potentially increase the frequency and intensity of catastrophic mass wasting events (Huggel et al., 2012). Accurate investigation and monitoring of mountain processes are therefore considered among the most important current activities in the field of environmental sciences, as they allow to detect indications and/or precursors of changes potentially impacting ecosystems at regional and global scales (Mark and Fernández, 2017).

The European Alps (hereafter referred to as Alps) are a key region of interest, because of their natural, cultural and economic resources for about 14 million inhabitants, and their touristic value attracting approximately 120 million visitors every year (The Alpine Convention: <http://www.alpconv.org/>, last access: 25 September 2025). Large slope instabilities are widespread in the Alps and are characterised by different mechanisms depending on the lithological, structural and morpho-climatic context of the slopes concerned. Instabilities can creep slowly for thousands of years because of long-term damage accumulation processes and suddenly undergo a progressive evolution towards catastrophic collapses (Riva et al., 2018; Agliardi et al., 2019, 2020, 2022). Due to their complexity, often including both creep behaviour and hydro-mechanical coupling (Crosta et al., 2017), the evolution towards failure events is difficult to anticipate. Furthermore, large landslides are usually characterised by complex styles of activity associated with separate slope sectors, expressed through different kinematics and strain partitioning into morpho/structural features (Crippa et al., 2020, 2021).

Surface deformation is a key indicator to assess spatial and temporal changes of landslide phenomena. For this reason, high resolution monitoring data of surface deformation is essential to investigate and interpret changes of landslide kinematics (Crosta et al., 2017; Loew et al., 2017). An important technique to monitor slope displacement is spaceborne Differential Synthetic Aperture Radar Interferometry (DInSAR, Wasowski and Bovenga, 2014). Since 2014, the availability of the ESA Copernicus Sentinel-1 mission has enhanced our capability to detect and monitor surface displacements also in landslide scenarios (Torres et al., 2012). Recently, an increasing number of authors have shown the potential for regional and national scale landslide monitoring by using different approaches exploiting the Sentinel-1 dataset (Barra et al., 2017; Manconi et al., 2018; Dehls et al., 2019; Raspini et al., 2018; Jones et al., 2023). Some intrinsic limitations of the Sentinel-1 system, however, hinder the nominal performance of standard and advanced spaceborne DInSAR methods in alpine settings, especially in cases where continuous monitoring is necessary to protect important economic and societal assets. The main limitations are related to (i) no information in areas affected by layover/shadow (Cigna et al., 2014), (ii) reduced or complete loss of displacement information in densely vegetated areas, as well as during periods of snow coverage, due to the scarce correlation of C-band SAR (wavelength ~ 5.6 cm), (iii) scarce or null capability of detecting displacements occurring

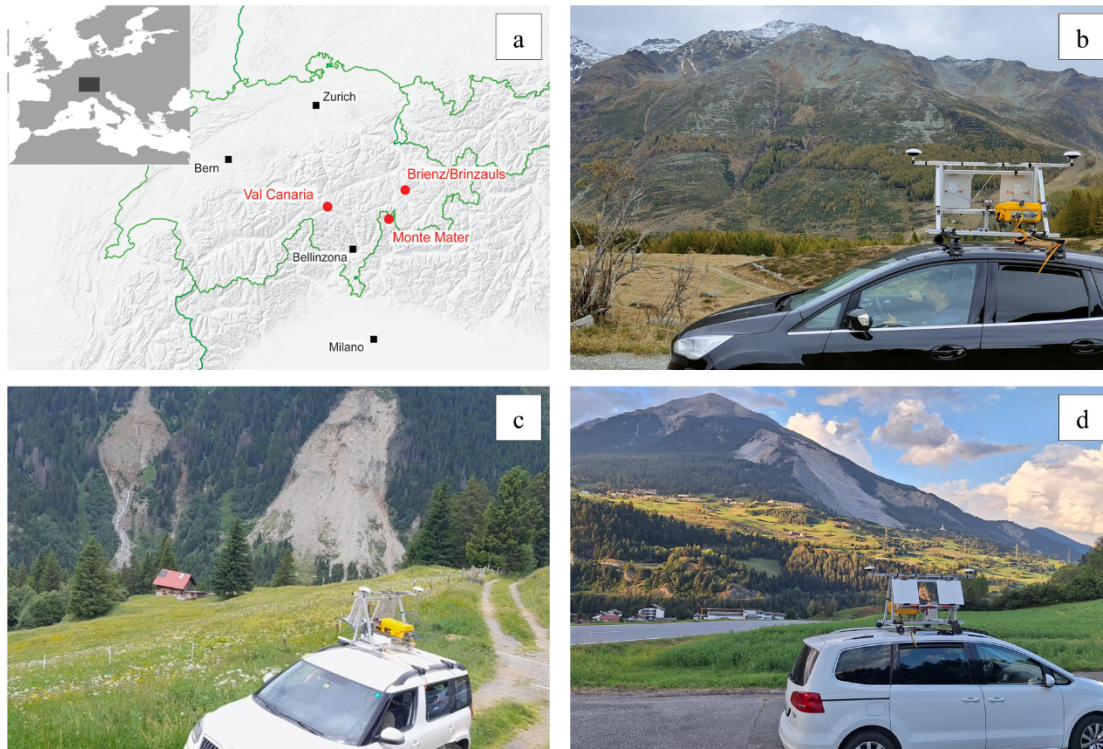
approximately along the satellite track direction (i.e., almost north-south direction), and (iv) inadequate revisit time of 6/12-days to follow the exponential evolution of surface displacements occurring prior to failure, due to phase ambiguity problems (Moretto et al., 2017; Manconi et al., 2018, 2021).

65 Considering these limitations, the use of multiple datasets is envisaged for a more complete analysis of surface displacements of unstable slopes. In particular, L-band (wavelength ~ 23.5 cm) satellites have the potential to complement the high frequency systems and, under certain conditions, provide information in vegetated areas and for comparatively large displacement rates (e.g., Strozzi et al., 2005; García-Davalillo et al., 2013; Nishiguchi et al., 2017; Ambrosi et al., 2018). ALOS-2 PALSAR-2 and SAOCOM-1 are currently two L-band satellite platforms with a good catalogue of data available
70 over the Alps, however, the acquisition strategy is not systematic. In the case of specific slope instabilities prone to catastrophic failure and that are potentially dangerous for human activities and infrastructure, ground-based methods are the sole alternative to set up a continuous monitoring and early warning system. While *in-situ* instrumentation such as robotized total stations, GNSS, cameras and extensometers are prominently considered for such applications, terrestrial radar remote sensing systems are becoming increasingly popular for early warning and emergency scenarios (Casagli et al., 2010; Crosta et al., 2014, 2017; Caduff et al., 2014; Caduff and Strozzi, 2017; Walter et al., 2019). Current terrestrial radar systems are
75 bound to relatively high frequencies, with many of them operating at Ku-band. A few years ago, GAMMA developed a new L-band SAR instrument (the GAMMA L-band SAR, GLSAR) and performed campaigns that demonstrate repeat-pass SAR interferometry for ground motion retrieval from agile platforms (Frey et al, 2019, 2021). Meanwhile, an updated and more compact version has been developed, which can also be mounted on multicopter drones (Frey et al., 2025).

80 The recent launch of the NISAR SAR satellite mission (NASA-ISRO) in July 2025 and the planned Copernicus Sentinel Expansion ROSE-L SAR satellite mission (ESA) will change the perspective on the use of L-band DInSAR products in landslide application, by providing free SAR imagery with a frequent, regular global coverage. Nevertheless, the full potential of these data in the mapping, activity characterisation, monitoring and failure forecast of alpine rock slope instabilities has still not been quantified in detail, hindering the preparedness of future users to consider such datasets. In this
85 work we perform a detailed analysis of ALOS-2 PALSAR-2 and SAOCOM-1 spaceborne and terrestrial GLSAR data to assess their potential for detecting and monitoring landslide occurrences in the Alps and provide an outlook for future L-band missions and instruments based on our findings.

2 Study Areas

We explore three different scenarios of large slope instabilities in the Alps that threaten important economic and societal
90 assets, with the support of the authorities responsible for landslide monitoring programmes in the region of Lombardia (Italy), Canton Ticino (Switzerland) and Canton Grisons (Switzerland) (Fig. 1).



95 **Figure 1:** (a) Spatial context of the three selected study sites in the European Alps. Val Canaria (Canton Ticino) and Brienzi/Brinzauls (Canton Grison) are located in Switzerland, Monte Mater (Lombardia) is located in northern Italy. (b) The Monte Mater rock slope instability in Italy with the GAMMA L-band SAR (GLSAR) instrument mounted on a car in the foreground. (c) The Rütan dei Sass landslide in Val Canaria with the car-borne GLSAR instrument in the foreground (cf. Section 3.2). Note that the measuring antennas do not face towards the landslide for illustration purposes only. (d) The Brienzi/Brinzauls landslide complex with the car-borne GLSAR instrument in the foreground. Note that the measuring antennas face towards the viewer for illustration purposes only.

100 2.1 Monte Mater, Lombardia, Italy

The Monte Mater rock slope instability is located in Valle Spluga (north-western Lombardia, Italy) and affects a 1300-metre-high glacial valley flank over an approximate area of 3 km² (Fig. 1b). The slope poses a threat to the village of Madesimo and the ski resort, which represents a significant asset in the area. The area was recognised as active based on inspection of Persistent Scatterer Interferometry (PSI) data and then monitored since 2011 by ARPA Lombardia (i.e., the public agency in charge of landslide monitoring and early warning in Lombardia) through periodical Ku-band Ground-Based (GB) DInSAR and GNSS measurements. Since 2017, the area was investigated in detail by combining *in-situ* geological, geomorphological and monitoring data and targeted DInSAR products derived from C-band Sentinel-1 imagery (Crippa et al., 2020). The study allowed identifying the nature and rates of the processes underlying the observed movements and provided a conceptual model to support the interpretation of the monitoring data. An active deep-seated gravitational slope deformation (DSGSD) affects the entire slope from toe to ridge, with translational global kinematics and estimated displacement rates of <3 cm/yr.

Two large landslides are nested within the DSGSD with partly rotational movements, strongly damaged rock masses and seasonally variable rates of up to 3-5 mm/month. In the upper part, scree and periglacial deposits move at faster rates up to 10 mm/month.

115 Monte Mater is an emblematic example of a challenging slope instability for all DInSAR techniques, due to its heterogeneity, variable rates of motion (in space and time) and environmental setting. The common loss of signal coherence in vegetated slope sectors, satellite line-of-sight (LOS) effects on the detection of some components of slow movements, atmospheric effects limiting the number of usable interferograms to analyse temporal trends of displacements and the difficulty of capturing processes occurring at different rates all contribute to the challenges of performing a detailed analysis of the local slope kinematics to support risk management for this class of phenomena. To discriminate processes
120 characterised by different mechanisms and displacement rates, reconstruct segmented sectors and identify temporal trends (linear, seasonal and non-linear), we complement existing interferometric datasets from C-band spaceborne and Ku-band GB SAR platforms with spaceborne and car-borne L-band measurements.

2.2 Val Canaria, Ticino, Switzerland

In Val Canaria, located near Airolo in the upper Leventina valley (Canton Ticino, Switzerland), several landslides affect
125 weak, steeply northwest-dipping metasediments of the Gotthard massif cover, which comprise mainly carbonate and gypsum rocks and display karst phenomena (Seno and Thüring, 2006). The landslides consist of failed and disrupted rock masses on both sides of the valley, with a total involved volume of about 80 million m³. This area is very close to the Gotthard transit corridor, an important infrastructure for road and rail connections between central and southern Europe. In the last twenty years, two main instabilities have been observed, which are well documented by geodetic measurements. These include the
130 rockfall of Ce on the left flank and the rockslide of Rütan dei Sass on the right flank. Both phenomena threaten to dam the Garegna torrent, with the consequent formation of flood waves endangering the cantonal road, highway and railway located in the valley bottom.

Monitoring data for Val Canaria's left flank, following the collapse on 27 October 2009 downstream of Ce that involved about 360'000 m³, include terrestrial DInSAR, GNSS, total station measurements, and terrestrial laser scanning data. Actual
135 displacements are on the order of 10 cm/a. On the right flank of Val Canaria, the active deep rockslide of the Rütan dei Sass involves a total volume of 10 Mm³ (Fig. 1c). This landslide has been monitored since 1992 with geodetic GNSS and theodolite measurements, which show displacements mainly on the order of few cm/yr. During landslide acceleration following heavy rainfall periods, however, a maximum displacement of 50 cm to 80 cm per day can be observed. Generally, the landslide has a retrogressive tendency, with the accumulation of numerous rockfalls that occur at the main scarp since
140 2015 and progressively cover the front of the rockslide. The last of these events occurred in 2018, involving a total volume of 170'000 m³. Currently, numerous tension cracks are present in the detachment zone, defining an estimated unstable

volume of about 8'000 m³ to 10'000 m³. In 2023 to 2024, the geodetic measurements recorded annual displacements between 20 and 30 cm that highlight the active state.

2.3 Brienz/Brinzauls, Grisons, Switzerland

145 A large landslide complex affects the southern flank of Piz Linard (located in Canton of Grisons, Switzerland) and involves
the village of Brienz/Brinzauls, hereafter referred to as Brienz (Fig. 1d). This site is extremely relevant in terms of landslide
hazard and risk potential, because it affects an entire slope with direct consequences for the Brienz village, its population and
the transportation infrastructures. The landslide recently experienced a rapid acceleration of surface displacements leading to
the evacuation of the village at the beginning of May 2023 and a slope failure event with a mass wasting of about 1.2 Mm³
150 on June 15th, 2023. In July 2023, the population of Brienz returned to their homes but were evacuated again in November
2024. To date, the village remains closed off to the population and visitors.

An extensive *in-situ* monitoring network has been progressively developed to measure surface deformation of the Brienz
landslide complex. This includes a combination of periodic surveys and permanent GNSS stations with a total of more than
80 points, as well as tens of reflectors measured with a robotized total station in operation since 2011 (Manconi et al., 2024).

155 In addition, surface displacements are also continuously monitored with a permanently installed GB radar (Loew et al.,
2024). Surface velocities on the order of several m/a are currently reported. These conditions are not ideal for the use of
DInSAR, especially for systematic monitoring applications. We investigate this site in detail with particular focus on the
accurate identification of surface velocities affecting the different landslide subdomains, also in areas affected by rapid
accelerations, the analysis of the surface velocities in 3D, by combining different SAR data sources (satellite and terrestrial),
160 and the identification of strategies for the systematic use of L-band DInSAR products within the current monitoring
programme.

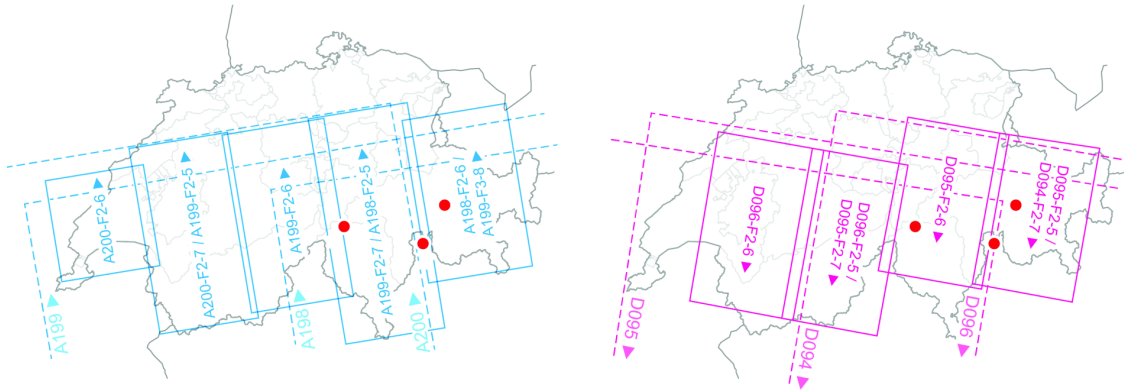
3 Datasets

We investigate landslide kinematics at the three study sites using standard and advanced DInSAR techniques based on
spaceborne L-band datasets from ALOS-2 PALSAR-2 and SAOCOM-1 satellites as well as terrestrial L-band measurements
165 acquired by the car-borne GLSAR. At the time of the results generation, spaceborne L-band data was available until the end
of 2023 for ALOS-2 PALSAR-2 and the end of 2024 for SAOCOM-1. A minimum of two car-borne GLSAR measurements
exist for each site, with more than 10 measurements available for Brienz.

3.1 ALOS-2 PALSAR-2

ALOS-2 PALSAR-2 provides L-band data with variable revisit times in Stripmap and ScanSAR mode, both of which are
170 used for this work (Strozzi et al., 2022). Considering the ALOS-2 PALSAR-2 archive over Switzerland, suitable Stripmap

mode data stacks in ascending orbit were systematically acquired over most of the Alps, while only specific regions (e.g., including the Brienz study site) were covered systematically by descending orbits. ScanSAR data, however, was acquired in systematic data stacks in descending orbit over the entire Swiss Alps and neighbouring regions, and only sporadically in ascending orbit. For this reason, in this work we combine snow-free Stripmap and ScanSAR data in DInSAR analyses for landslide velocity mapping. The three key study regions are each covered by ascending and descending ALOS-2 PALSAR-2 orbits as shown in Fig. 2. Table 1 lists the orbits and number of acquisitions used for this study. Lists of the ALOS-2 PALSAR-2 acquisitions used in the analysis at each study site are provided in Appendix A1, B1, C1 and C2.



180 **Figure 2: Ascending (left) and descending (right) ALOS-2 PALSAR-2 orbits covering the three study regions, which are marked by red dots. ScanSAR orbits are outlined with dashed lines, Stripmap orbits with continuous lines.**

Table 1: Processed ALOS-2 PALSAR-2 data per study site. Orbits in brackets are available as illustrated in Fig. 2 but are not used in the analysis.

Study Site	Ascending	Processed data	Descending	Processed data
Monte Mater, IT	A199-F2-7 (A198-F2-5)	24 snow-free acquisitions from 2014-2023	D095-F2-5 (D094-F2-7)	37 snow-free acquisitions from 2015-2023
Val Canaria, CH	A199-F2-7 (A198-F2-5)	24 snow-free acquisitions from 2014-2023	D095-F2-6	23 snow-free acquisitions from 2015-2023
Brienz, CH	A199-F3-8 (A198-F2-6)	28 snow-free acquisitions from 2016-2023	D095-F2-5 (D094-F2-7)	40 snow-free acquisitions from 2015-2023

3.2 SAOCOM-1

185 SAOCOM-1 provides L-band single, dual or quad polarization SAR data with a revisit time of 8/16 days (SAOCOM: <https://earth.esa.int/eogateway/missions/saocom/>, last accessed 27 February 2025). We use Stripmap mode data, which are regularly acquired over the Alps. Individual SAOCOM-1 Stripmap scenes are approximately 70 km wide and data acquisitions started at the end of 2020, however, regular acquisitions every 16 or 32 days along ascending and descending orbits are only available since 2023. SAOCOM-1 has a much larger orbital tube than other current satellite SAR missions,

such as ALOS-2 PALSAR-2 and Sentinel-1, with perpendicular baselines partly exceeding 1000 m. In addition, inaccuracies in the annotated state vectors of the input single look complex images introduce errors in the baseline calculation, which in turn leave a residual phase contribution in the interferogram. Despite this, very promising DInSAR results can be achieved with SAOCOM-1 Stripmap data (Strozzi et al., 2024; De Luca et al., 2025). The three key study regions are each covered by an ascending and descending SAOCOM-1 orbit as shown in Fig. 3. Table 2 lists the orbits and number of acquisitions used for this study. Lists of the SAOCOM-1 acquisitions used in the analysis at each study site are provided in Appendix A1, B1, C1 and C2.

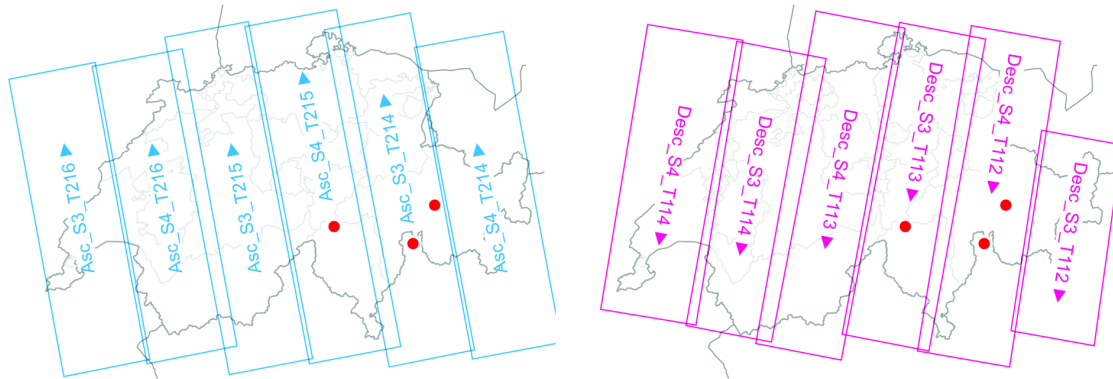


Figure 3: Ascending (left) and descending (right) SAOCOM-1 orbits covering the three study regions, which are marked by red dots.

Table 2: Processed SAOCOM-1 data per study site.

Study Site	Ascending	Available data	Descending	Available data
Monte Mater, IT	S3 T214 Frame 528	37 snow-free acquisitions from 2021-2024	S4 T112 Frame 74	31 snow-free acquisitions from 2021-2024
Val Canaria, CH	S4 T215 Frame 528	27 snow-free acquisitions from 2020-2024	S3 T113 Frame 73	21 snow-free acquisitions from 2022-2024
Brienz, CH	S3 T214 Frame 528	42 snow-free acquisitions from 2021-2024	S4 T112 Frame 73	21 snow-free acquisitions from 2022-2024

200 3.3 GAMMA L-band SAR instrument

The GAMMA L-band SAR (GLSAR) instrument can be mounted on terrestrial rail-, car-, or airborne platforms and acquires radar images at L-band with a maximum bandwidth of 200 MHz within 1.2-1.4 GHz (Frey et al., 2019, 2021, 2023, 2024, 2025). For terrestrial analyses at the three study sites, the GLSAR was used in the car-borne configuration to perform repeat-pass interferometry to retrieve LOS displacements on the slope of interest affected by ground motion (Fig. 1b-1d). Only 100 MHz bandwidth around a centre frequency of 1.325 GHz were used due to regulatory restrictions. The car-borne SAR system setup includes a GNSS-aided inertial navigation system (INS) and a local ad-hoc GNSS reference station to precisely

determine the sensor trajectory during the SAR data acquisition. Using this 3-D sensor trajectory and a digital elevation model of the scene, geocoded single-look complex (GEOSLC) SAR images are produced with a parallelized time-domain back-projection (TDBP) algorithm (Frey et al. 2009) implemented for NVIDIA GPUs (Frey et al. 2014). This ensures that the SAR image is georeferenced with high geodetic accuracy and that topography-corrected differential interferograms can readily be computed from the GEOSLCs. Two measurements each were acquired at Monte Mater on 14 October 2024 and 10 June 2025 and at Val Canaria on 3 July and 23 September 2024. More than 10 measurements are available at Brienz between 2020 and 2024 (Appendix C3), of which we consider those acquired on 16th and 23rd September 2023.

4 Methods

We employ standard and advanced DInSAR methods for ground motion detection and landslide velocity mapping based on the input data as outlined in Section 3. Multiple studies have investigated the potential and limitations of these techniques, using mainly C-band Sentinel-1 data (Darvishi et al., 2018; Crippa et al., 2020; Jones et al., 2021, 2023; Manconi et al., 2018, 2024; Strozzi et al., 2018, 2020). In the following work, we analyse the potential of L-band data for DInSAR uses and aim to complement results from higher frequency data. Specifically, we mainly apply PSI to detect ground motion from satellite data on a scale of few centimetres per year (Ferretti et al., 2001; Wegmüller et al., 2003; Bianchini et al., 2013; Cigna et al., 2013). In regions of faster movement and poor coverage with PSI, additional methods such as standard DInSAR, stacking and Small Baseline Subset (SBAS) are applied (Lundgren et al., 2001; Berardino et al., 2002; Schmidt and Bürgmann, 2003; Werner et al., 2012). Where available, we combine data from both ascending and descending orbits for maximum spatial coverage of the three key study regions.

4.1 Persistent Scatterer Interferometry

PSI allows surface movement to be detected at millimetre scale accuracy. The method is based on the selection of temporally coherent pixels termed Persistent Scatterers (PS) to measure surface displacement in a stack of unwrapped differential interferograms (Ferretti et al., 2001; Wegmüller et al., 2003; Colesanti and Wasowski, 2006;). Resulting displacement maps allow the extraction of time series for individual points at high resolution and accuracy (Crosetto et al., 2009). As PS are typically smaller than the resolution cell size, PSI is a suitable method also for sensors with large spatial baselines such as SAOCOM-1. We perform PSI according to the Interferometric Point Target Analysis (IPTA) algorithm, which selects PS in two approaches (Wegmüller et al., 2003). A first selection is performed based on spectral diversity, followed by a second selection based on the intensity temporal variability. This approach exploits both the spectral and intensity properties of the input images. The resulting PS are merged into a single comprehensive candidate list and used in a subsequent point target quality analysis in differential interferograms, which filters PS based on selected spatial and/or temporal coherence thresholds. The displacement at these point locations is then derived in an iterative two-dimensional regression process and in reference to a point of known or assumed zero displacement. This reference point is selected manually in areas considered

to be stable. The final geocoded displacement map allows extraction of velocity time series for individual points to show the temporal evolution of displacement in specific areas of interest.

240 **4.2 Standard DInSAR**

Standard DInSAR encompasses the generation of differential interferograms by extracting the difference in phase information of two SAR images covering largely the same area and acquired at different times (Bürgmann et al., 2000). By removal of the topographic phase through use of an external DEM, a value of surface displacement along the satellite's LOS is derived in the resulting image (Wasowski and Bovenga, 2014). Removal of atmospheric disturbances is performed
245 empirically with a high pass filter which attenuates the low-frequency elevation-dependent atmospheric components and preserves the high-frequency components of the landslide motion. Absolute displacement values in the LOS direction are extracted through phase unwrapping of the modulo 2π ambiguity, typically by applying a minimum cost flow algorithm (Costantini, 1998; Werner and Wegmüller, 2002). A reference point of known or assumed zero displacement is used for processing and typically chosen in stable rock outside the identified moving area. Unwrapping procedures can introduce
250 errors particularly in regions of high topographic relief, such as the Alps (Berardino et al., 2002; Darvishi et al., 2018). Areas of low coherence due to fast movement or vegetation/snow cover may not be successfully unwrapped in differential interferograms and can introduce an over- or underestimation of surface movement in subsequent displacement maps. By choosing specific, small areas of interest for unwrapping procedures, associated errors can be minimised. While unwrapped differential interferograms are used as the input for advanced DInSAR methods, we also considered wrapped differential
255 interferograms for the inspection of areas for which PSI does not provide information, e.g. in gaps caused by interferometric decorrelation due to large motion. Wrapped differential interferograms show the derived surface displacement as coloured fringe cycles on the order of a 2π ambiguity, relating to half the satellite wavelength (i.e., 11.8 cm for L-band sensors).

4.3 DInSAR Stacking

Stacking approaches combine unwrapped differential interferograms of variable temporal and spatial baselines over a
260 specific period of interest to obtain average displacement maps within that time frame (Lundgren et al., 2001). Constraints on the spatial baselines are not strictly necessary but can be applied to increase overall coherence. Stacking is particularly suitable for seasonal and annual deformation analyses and typically requires the exclusion of winter images in the Alps due to extensive snow cover and resulting loss of coherence. This approach ensures sufficiently high coherence but introduces data gaps in the resulting displacement time series. Stacking is a popular complementary method to DInSAR techniques that
265 detect millimetre-scale movements such as PSI and has been applied in an extensive study of C-band Sentinel-1 data at Brienz (Manconi et al., 2024).

4.4 Small Baseline Subset

SBAS combines unwrapped differential interferograms of small spatial and temporal baselines to ensure that spatial coherence is maintained (Berardino et al., 2002; Wasowski and Bovenga, 2014). This is particularly critical for sensors with large and variable spatial baselines such as SAOCOM-1. SBAS is particularly suitable for seasonal displacement analyses but can also be implemented to study longer-term displacement. As for PSI, reference points of known or assumed zero displacement are defined for each of the three study regions. Displacement values are extracted from the stacks of unwrapped differential interferograms in reference to this (stable) point and displayed in displacement maps and time series, showing the progressive change in LOS surface deformation for the assessed time period. Masking is applied to remove areas of low coherence, e.g., due to snow/vegetation cover or very fast movements that exceed the phase aliasing threshold of $\lambda/4$. Large-scale spatial filtering removes errors related to large-scale atmospheric effects.

4.5 GAMMA L-band SAR instrument

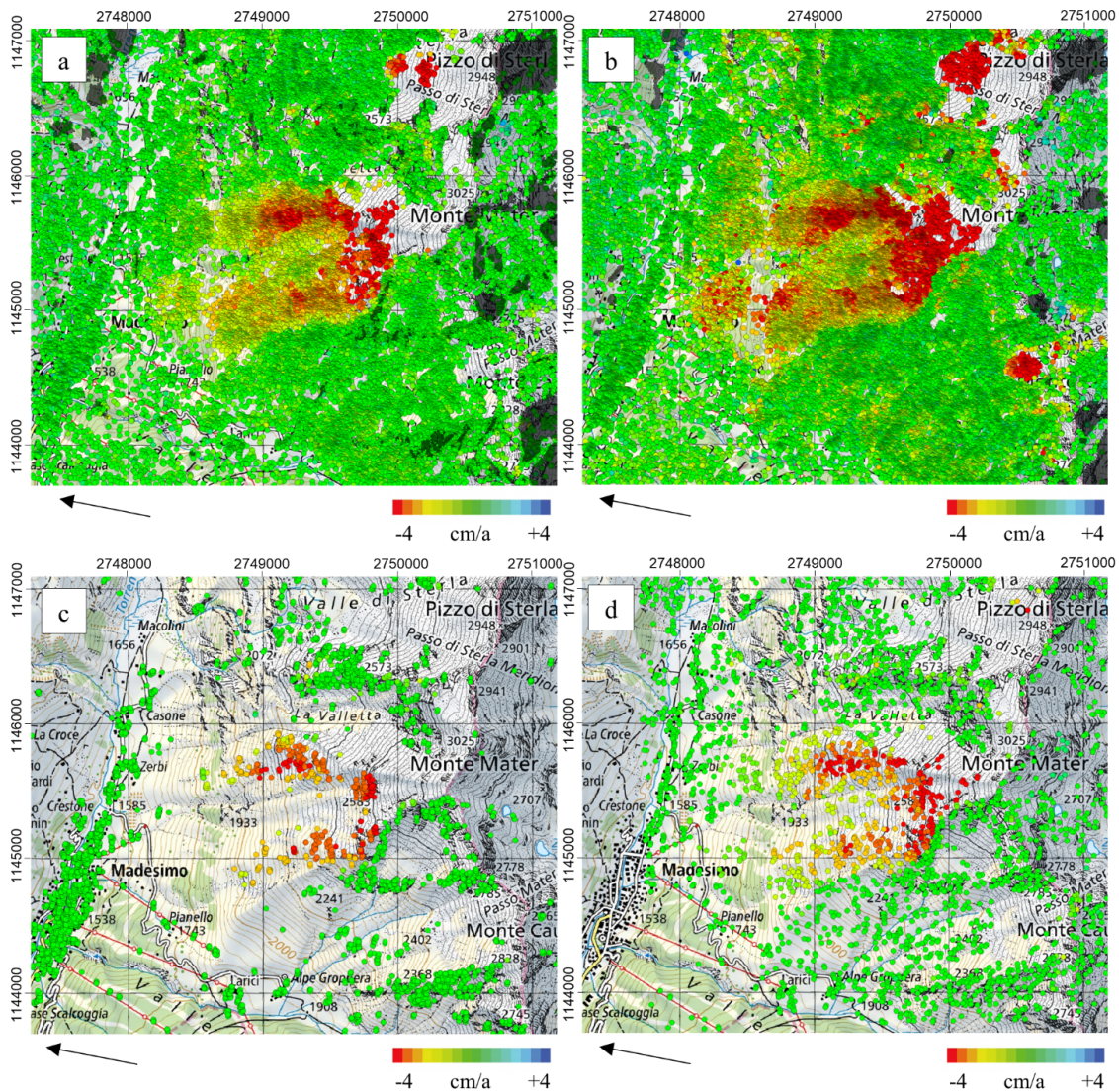
Terrestrial DInSAR data is acquired through use of the car-borne GLSAR system (see Section 3.3). It is employed locally on one or several days and measures surface displacement within a time interval of a few hours, days or longer time frames, depending on the velocity of the observed mass movement. A minimum of two independent acquisitions is required to compute topography-corrected differential interferograms. As for satellite data, DInSAR data displays ground movement in wrapped format and displacement maps are generated after phase unwrapping. A reference point of known or assumed zero displacement is used for processing and typically chosen in stable rock outside the identified moving area. Removal of atmospheric disturbances is performed empirically with a high pass filter which attenuates the low-frequency atmospheric components (mainly range and elevation dependent) and preserves the high-frequency components of the landslide motion.

5 Results

5.1 Monte Mater, Lombardia, Italy

The PSI results at the Monte Mater study site for descending ALOS-2 PALSAR-2 and SAOCOM-1 data are illustrated in Fig. 4. The landslide is clearly identified, with areas depicted as red when moving away from the satellite and as blue when moving towards the satellite. The point density is slightly higher in SAOCOM-1, particularly in the upper parts of the slope, due to its higher temporal sampling rate of acquisitions within a shorter observation time interval. However, SAOCOM-1 also shows higher levels of noise in the surrounding stable areas. This is due to the shorter observation time interval of SAOCOM-1 (2021-2024) compared to ALOS-2 PALSAR-2 (2015-2023). The Monte Mater site is very well illuminated by satellite data of the descending orbit. For the ascending orbit, the satellite LOS and the direction of motion are approaching a perpendicular direction. Large areas are affected by layover, especially in SAOCOM-1 images, and motion is observed

towards the satellite. The PSI results at the Monte Mater study site for ascending ALOS-2 PALSAR-2 and SAOCOM-1 data are illustrated in Appendix A2.

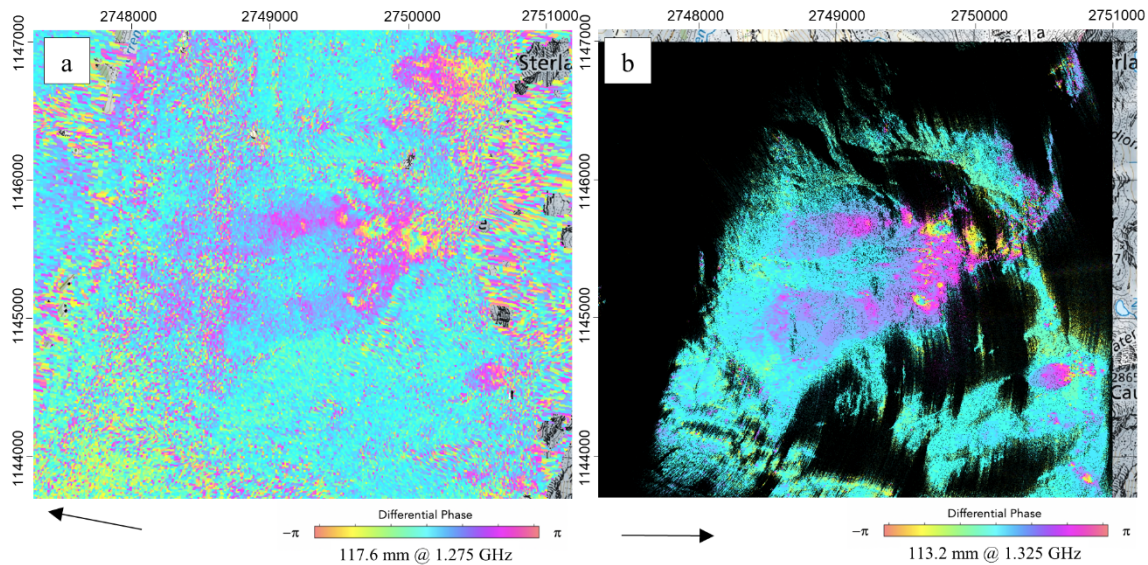


300 **Figure 4: Results of PSI analysis at Monte Mater with ALOS-2 PALSAR-2 for the time period 2015–2023 (a), SAOCOM-1 for the time period 2021–2024 (b), Sentinel-1 for the time period 2018–2022 from the European Ground Motion Service (c; EGMS; <https://egms.land.copernicus.eu/>; last accessed 22 April 2025) and SqueeSAR for the time period 2015–2017 (d; Crippa et al., 2021). The arrows indicate the satellite LOS direction for the descending orbit. Layover/shadow is shown in black. Image background is the Swiss National Map.**

C-band datasets using Sentinel-1 acquisitions at this site are available through the European Ground Motion Service (EGMS: 305 <https://egms.land.copernicus.eu/>; last accessed 22 April 2025) for the period 2018–2022 and from SqueeSAR for the period 2015–2017 (Crippa et al., 2021), and are illustrated in Fig. 4c and 4d. The high density of PS in L-band and their spatial

continuity, granted by a higher interferogram coherence in all the landcover conditions, allow a better delineation of the kinematic domains forming the Monte Mater landslide complex, especially the southern nested rockslide and the upper part of the DSGSD. L-band data also better illuminate the internal heterogeneity and toe activity of the nested landslides.

- 310 A wrapped differential interferogram for descending SAOCOM-1 data from 14 October 2024 to 11 June 2025 is shown in Fig. 5a and supports and complements the PSI results. Despite a temporal baseline of 240 days, very high coherence is maintained, allowing to retrieve spatially distributed information. This translates into the ability to capture different associated processes acting on different temporal rates. In particular, the L-band wrapped differential interferogram clearly indicates high displacement rates in the upper parts of the slope as visualised by multiple fringe cycles or even decorrelation.
- 315 Terrestrial measurements of the Monte Mater landslide were acquired on 14 October 2024 and 10 June 2025 with the car-borne GLSAR. Several repeat-pass SAR measurements were acquired on each day within a total time span of approximately one hour. The averaged interferogram is shown in Fig. 5b and indicates good overall coherence. The landslide is clearly identified and, as for descending SAOCOM-1 data from nearly the same period in Fig. 5a, indicates high displacement rates in the upper parts of the slope, but with higher spatial resolution and from a different look direction.



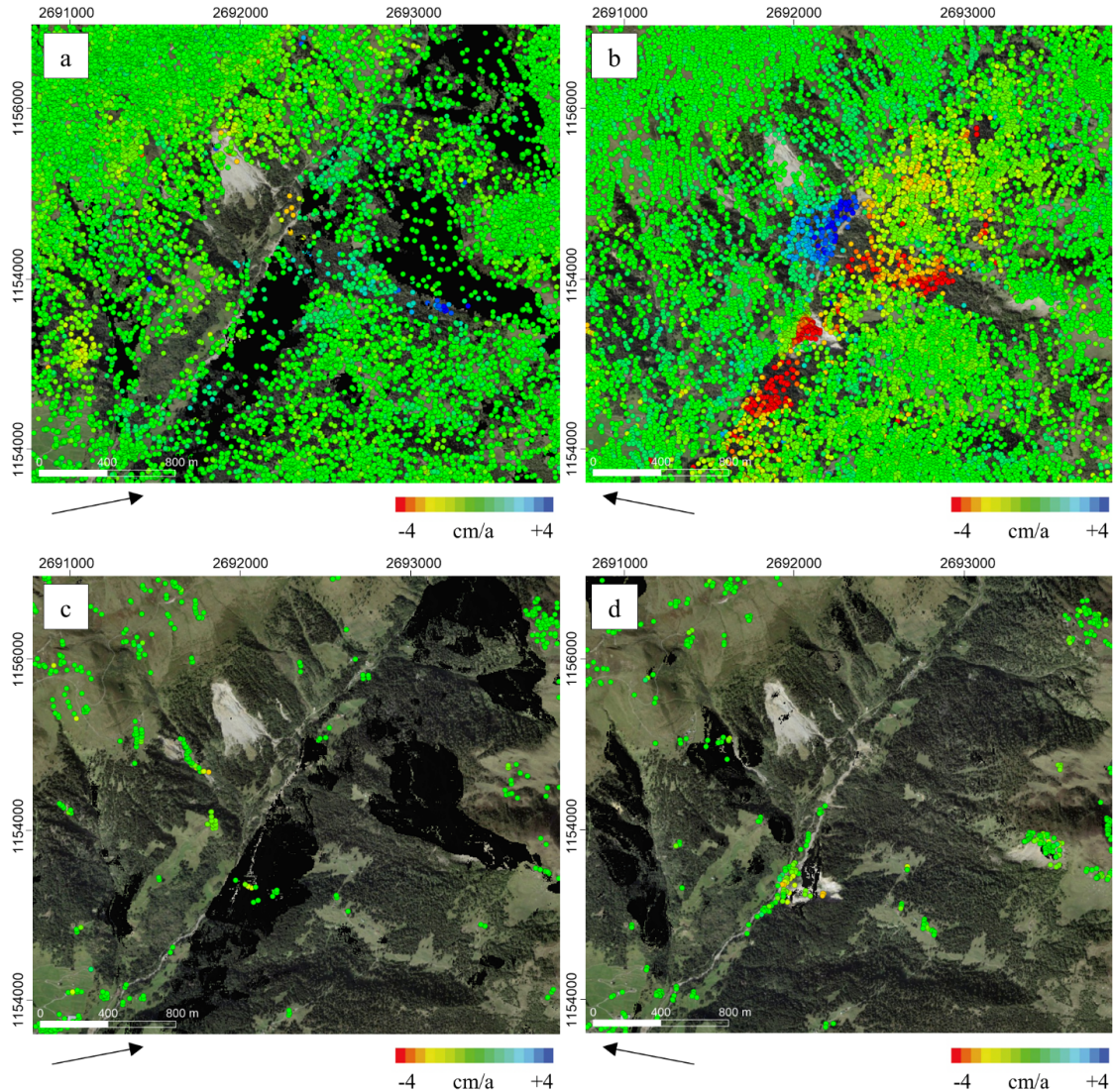
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Figure 5: Descending SAOCOM-1 interferogram over 14 October 2024 – 11 June 2025 (a) and car-borne GLSAR interferogram over 14 October 2024 – 10 June 2025 (b) at Monte Mater. The arrows indicate the satellite and terrestrial LOS directions. Layover/shadow is shown in transparent (a) and black (b). Image background is the Swiss National Map.

5.2 Val Canaria, Ticino, Switzerland

- 325 Satellite L-band PSI results at the Val Canaria study site are illustrated in Fig. 6a and 6b, with SAOCOM-1 overlain on ALOS-2 PALSAR-2, for ascending and descending orbits. The results show overall high point coverage with some data gaps remaining on both sides of the valley. SAOCOM-1 generally increases the point density compared to ALOS-2 PALSAR-2,

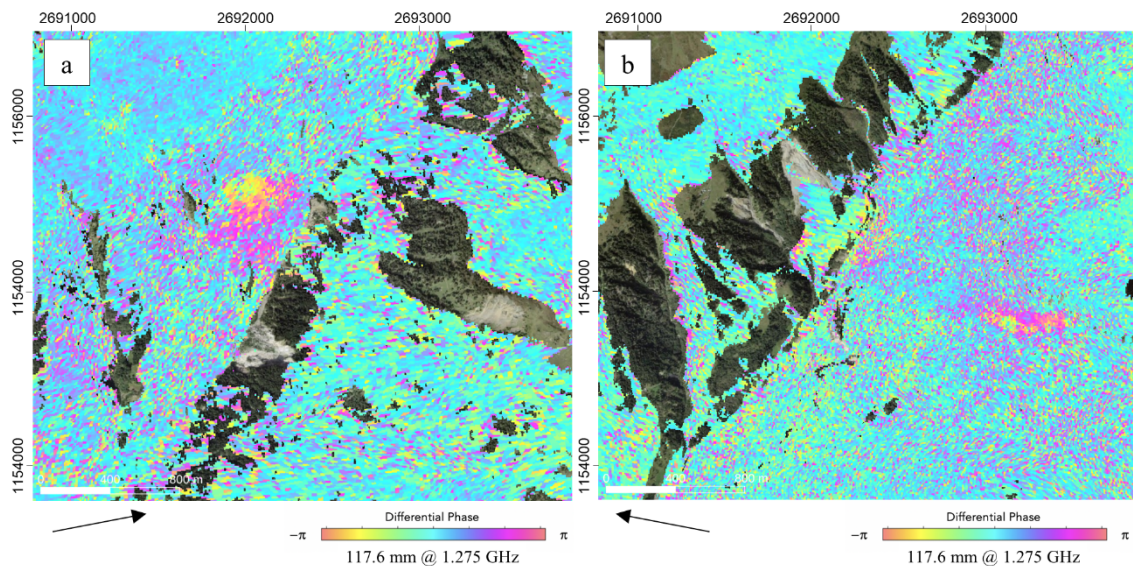
due to its higher temporal sampling rate of acquisitions within a shorter observation time interval. This is particularly apparent in descending orbit, where several moving areas are detected. PSI results for only ALOS-2 PALSAR-2 and SAOCOM-1 are provided in Appendix B.2 and B.3. Sentinel-1 C-band PSI results, available through Geoportale Ticino (<https://map.geo.ti.ch/>; last accessed 22 April 2025), are presented in Fig. 6c and 6d and provide very limited coverage. Especially in forested sectors the landslide detection and characterisation are scarce.



335 **Figure 6: Results of PSI analysis at Val Canaria with ascending SAOCOM-1 for the time period 2020–2024 overlain on ALOS-2 PALSAR-2 for the time period 2014–2023 (a), with descending SAOCOM-1 for the time period 2022–2024 overlain on ALOS-2 PALSAR-2 for the time period 2015–2023 (b) and with ascending (c) and descending (d) Sentinel-1 for the time period 2015–2019 (Geoportale Ticino: <https://map.geo.ti.ch/>; last accessed 22 April 2025). The arrows indicate the satellite LOS direction. Layover/shadow is shown in black. Image background is SWISSIMAGE.**

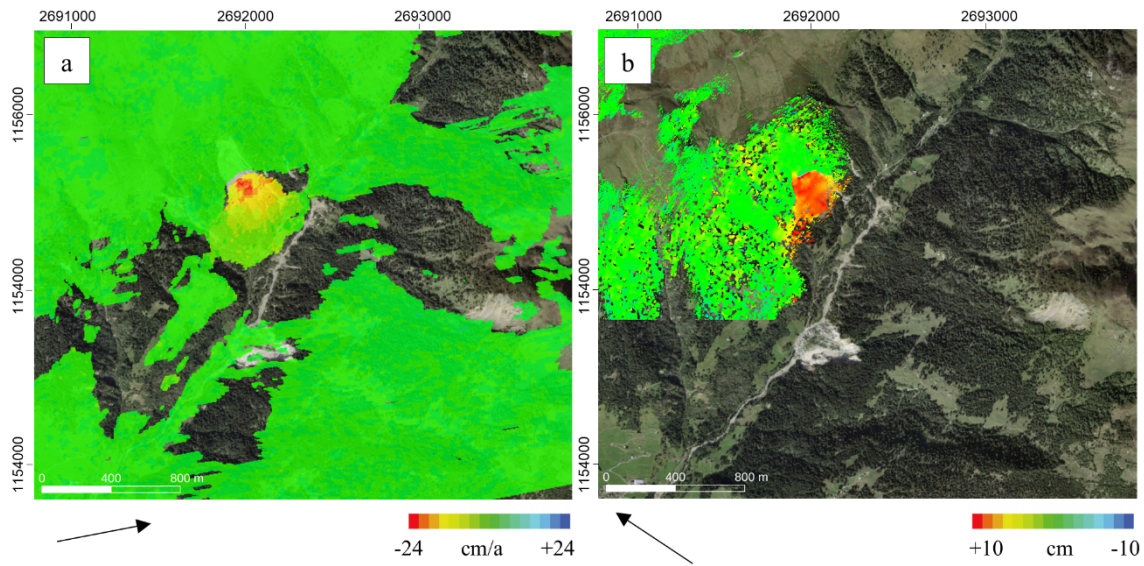
The main landslide of interest on the southeast-facing slope of the study site is only partly detected in descending orbit and
340 not at all in ascending orbit. This data gap and lack of displacement information likely indicate conditions of high velocities
that cannot be resolved with the PSI technique even at L-band, as the underlying orthophoto shows a scree slope that is
partly free of vegetation at this location. SAOCOM-1 wrapped differential interferograms as illustrated in Fig. 7 support the
PSI results and confirm high displacement rates on the affected slope, with a minimum of one fringe cycle per temporal
baseline.

345 To complement the PSI results and fill data gaps related to fast surface movements, alternative advanced DInSAR techniques
are considered. Ascending SAOCOM-1 data was used for SBAS with perpendicular baselines shorter than 750 m and
acquisition time intervals of less than 270 days to enable the detection of higher velocities while maintaining coherence. The
results highlight that part of the southeast-facing slope at Val Canaria moves with estimated velocities of >20 cm/a, clearly
exceeding the detectable threshold with PSI (Fig. 8a). A corresponding LOS displacement time series (Figure 13e discussed
350 in Section 6.2) indicates non-linear motion that appears to have accelerated in 2021 and 2024. Terrestrial measurements of
the southeast-facing slope were acquired on 3rd July and 23rd September 2024 with the car-borne GLSAR. The displacement
map is shown in Fig. 8b and indicates good overall coherence and velocities on the order of 10 cm in nearly three months.
The landslide is clearly detected with a sharp delineation of its upper extent, as also visible for the SAOCOM-1 SBAS
results of Fig. 8a.



355

Figure 7: Ascending SAOCOM-1 interferogram over 24 July – 12 October 2024 (a) and descending SAOCOM-1 interferogram over 25 August 2022 – 7 October 2023 (b) at Val Canaria. The arrows indicate the satellite terrestrial LOS directions. Areas of layover/shadow are shown as transparent. Image background is SWISSIMAGE.



360 **Figure 8: SBAS results from ascending SAOCOM-1 for the time period 2020–2024 (a) and car-borne GLSAR LOS displacement**
 365 **map between 3 July and 23 September 2024 (b). The arrows indicate the satellite and terrestrial LOS directions. Areas of**
 370 **layover/shadow are shown as transparent. Image background is SWISSIMAGE.**

5.3 Brienz/Brinzauls, Grisons, Switzerland

ALOS-2 PALSAR-2 PSI results at the Brienz study site are shown in Fig. 9a and 9b for ascending and descending orbits.

365 The point density is overall very high, with two large moving areas detected further northwest near Lantsch/Lenz as well as
 370 to the east of Brienz. The Brienz landslide itself, however, which includes the slopes above and surrounding the village of
 Brienz, remains a significant data gap. This is due to the very fast surface displacement observed here that cannot be
 resolved with PSI even at L-band. SAOCOM-1 PSI results do not provide additional information despite the satellite's
 higher temporal sampling rate, as shown in Appendix C.4. Overall, they are noisier and contain less information at higher
 altitudes compared to ALOS-2 PALSAR-2, due to the shorter total observation interval, lower data availability (e.g., when
 comparing with the Monte Mater study site) and use of winter data without snow cover over Brienz at low altitudes. C-Band
 cannot provide further information under forested areas and in rapidly decorrelating slopes, as shown with Sentinel-1 PSI
 results in Fig. 9c and 9d for ascending and descending orbits (Strozzi et al., 2022). Additionally, due to the near-polar flight
 path of both satellites, the detection sensitivity on the overall south-facing orientation of the slope is generally lower.

375 Due to the limitations of PSI for the detection of high displacement rates, alternative advanced DInSAR techniques such as
 stacking are required. SBAS is not possible here as both L-band satellite sensors do not provide a temporally continuous
 series of differential interferograms, due to partly very long spatial or temporal baselines as well as decorrelation in winter
 because of snow cover. Stacking enables the detection of fast displacements by considering several differential
 interferograms over a specified time frame to obtain an average displacement rate and reduce random noise. Differential

380 interferograms in 2024 from ascending SAOCOM-1 track S3 T214, with nearly consistent 16 days acquisitions, and descending ALOS-2 PALSAR-2 track 095-F2-5, with temporal baselines of mostly 14 days as well as some 28 days and longer intervals, are most suitable for this technique. Appendix C.2 lists the differential interferograms used in the analysis. The stacking results in Fig. 10 show that ascending SAOCOM-1 and descending ALOS-2 PALSAR-2 data give a good overview of the LOS displacement over the lower part of the Brienz landslide. Large data gaps remain in the upper parts of the slope, however, due to very high local movement rates as well as layover and shadowing effects.

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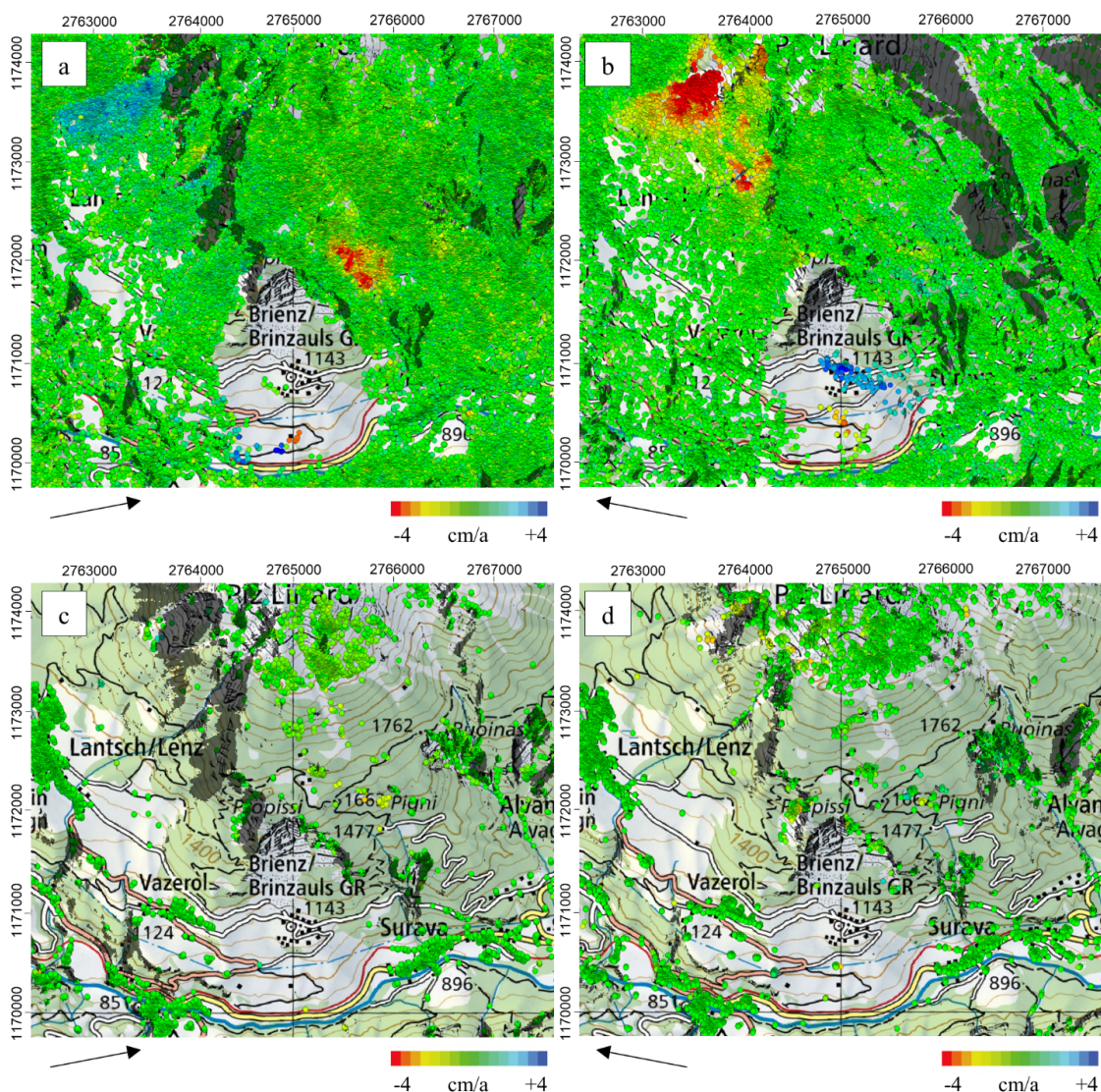


Figure 9: Results of PSI analysis at Brienz with ascending ALOS-2 PALSAR-2 for the time period 2019–2023 (a), descending ALOS-2 PALSAR-2 for the time period 2015–2023 (b) and with ascending (c) and descending (d) Sentinel-1 for the time period 2015–2021. The arrows indicate the satellite LOS direction. Layover/shadow is shown in black. Image background is the Swiss National Map.

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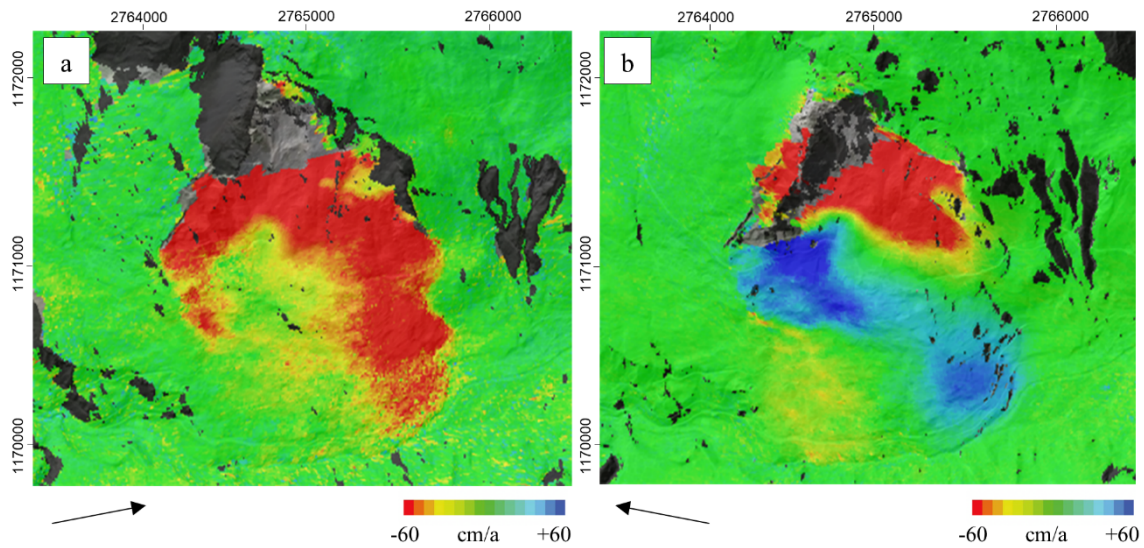
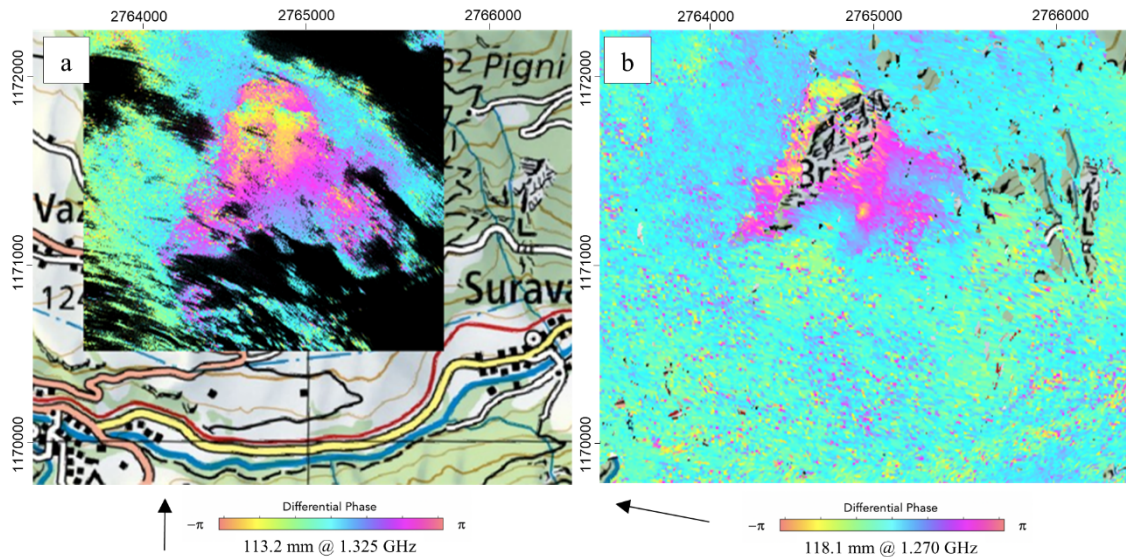


Figure 10: Results of DInSAR stacking analysis in 2024 at Brienz with ascending 16-day SAOCOM-1 (left) and descending 14/28-day ALOS-2 PALSAR-2 (right). Appendix C.2 lists the differential interferograms used in the analysis. The arrows indicate the satellite LOS direction. Layover/shadow is shown in black. Image background is SWISSIMAGE.

395 Terrestrial repeat-pass SAR measurements of the Brienz landslide were acquired >10 times between 2020 and 2024 with the
 car-borne GLSAR as listed in Appendix C.3. Considering the acquisitions from 16 and 23 September 2023, i.e. after the
 rockfall in June 2023, the resulting differential interferogram with a temporal baseline of one week clearly outlines the
 landslide and indicates different compartments of varying displacement rates within the moving area (Fig. 11a). Compared to
 a differential interferogram of two weeks obtained for descending ALOS-2 PALSAR-2 (Fig. 11b), the overall coherence of
 400 the GLSAR-derived interferogram is higher. The line of sight of the car-borne SAR acquisition aligns very well with the
 north-south orientation of the landslide's motion vector field, whereas this is not the case for the spaceborne acquisitions.
 Furthermore, in the satellite observations, areas at higher altitudes are obscured by layover or shadows, while visibility from
 the ground in this sector is very good. The GLSAR interferogram also corresponds very well with the results of digital image
 correlation (DIC) on high-resolution multi-temporal digital terrain models (DTM) created from airborne surveys along the
 405 north-south direction (Manconi et al., 2024). The GLSAR results thus enable the detection of coherent displacement in the
 fastest moving parts of the landslide at higher altitudes and clearly highlight the complementary potential of terrestrial L-
 band measurements to the available satellite L-band data. Both datasets show displacement of at least one fringe cycle during
 a temporal period of one and two weeks, respectively, confirming that movement rates exceed the threshold detectable with
 the PSI technique.



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Figure 11: Interferograms for car-borne GLSAR data acquired on 16 and 23 October 2023 (a) and descending ALOS-2 PALSAR-2 data acquired on 7 and 21 September 2023 (b). The arrows indicate the terrestrial and satellite LOS directions. Layover/shadow is shown in black. Image background is the Swiss National Map.

6 Discussion

415 6.1 L-band versus C-band

C-band datasets using Sentinel-1 acquisitions are available at the three test sites through the European Ground Motion Service (EGMS; <https://egms.land.copernicus.eu/>; last accessed 22 April 2025), Geoportale Ticino (<https://map.geo.ti.ch/>; last accessed 22 April 2025) and independent studies (Crippa et al., 2021; Strozzi et al., 2022). The comparison of C- and L-band datasets across the three alpine study sites highlights pronounced differences in displacement detection. Histograms of PS density were computed for all the considered datasets for forested and non-forested terrain and are presented in Fig. 12. Size and percentage of forested and non-forested terrain per area (Table 3 and Appendices A3, B4 and C5) are based on the CORINE Land Cover (CLC, <https://land.copernicus.eu/en/products/corine-land-cover/>, last accessed 31 March 2026). As also apparent in Fig. 1, forest cover is predominant in Val Canaria (63%), more limited in Brienz/Brinzauls (31%), and sparse in the alpine environment of Monte Mater (8%). The histograms of PS density confirm that L-band acquisitions consistently provide much higher PS densities, in some cases up to twenty times greater than C-band, depending on the satellite platform, orbit direction, and number of images available in the multi-temporal stacks.

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At Monte Mater in northern Italy, the contrast is particularly striking. In unforested terrain, the PS density of L-band products exceeds that of C-band by a factor of up to 20, while in forested areas the density remains 10–15 times higher. This increase in point density at L-band allows the moving area of the DSGSD and its internal kinematic compartments to be

430 resolved with a spatial continuity that is impossible to achieve with C-band products, which display only scattered clusters of points.

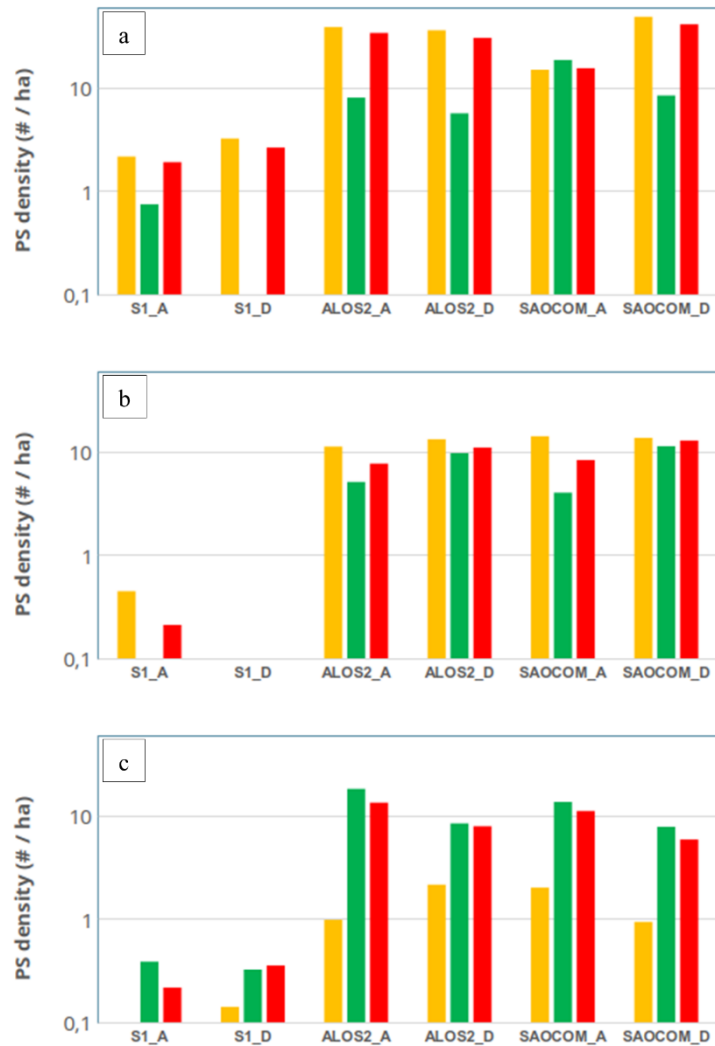
In Val Canaria, the limitations of C-band are even more apparent. Sentinel-1 analyses yield very low point densities overall and no solution at all on vegetated slopes, leaving the mapped landslides essentially undetected. In contrast, L-band acquisitions from ALOS-2 PALSAR-2 and SAOCOM-1 provide dense PS coverage throughout the SAR-visible portions of the slope, with average densities approximately five times higher than those of C-band. The improvement is consistent across both forested and non-forested areas, demonstrating that L-band maintains coherence where C-band fails. In addition, the higher temporal sampling rate of SAOCOM-1 compared to ALOS-2 translates directly into a denser point distribution and better definition of slope sectors with more rapid motion.

The Brienz/Brinzauls landslide illustrates both the advantages and the limits of the approach. C-band analyses show sparse coverage and the large moving areas to the north-west near Lantsch/Lenz and east of Brienz remain undetected. L-band datasets, by contrast, clearly reveal these zones with a significantly higher PS density, especially in forested areas. The reason of lower point densities in non-forested areas is because these areas are also the ones in the central body of the landslide experiencing the highest velocities in the period of observation. Where displacements reached very high values, coherence is lost in both frequency bands and no PSI points are retrieved. This reflects a fundamental limit of the PSI technique rather than a wavelength effect: once velocities exceed a certain threshold, neither C- nor L-band can provide stable measurements. Due to the limitations of PSI for the detection of high displacement rates, alternative techniques such as standard DInSAR or stacking can be employed. Here again, with its superior coherence across all temporal baselines, L-band can better retrieve spatially distributed information and capture various processes acting at different velocities.

Table 3: Size and percentage of non-forested/forested terrain per area in which the PS density is calculated for each test site.

Site	Total area [km ²]	Non-forested area [km ² / %]	Forested area [km ² / %]
Monte Mater	3.21	2.94 / 92%	0.27 / 8%
Val Canaria	3.88	1.45 / 37%	2.44 / 63%
Brienz/Brinzauls	3.12	2.15 / 69%	0.97 / 31%

450



455 **Figure 12: PS densities for products derived from Sentinel-1 (S1), ALOS-2 PALSAR-2 (ALOS2) and SAOCOM-1 (SAOCOM) in ascending (A) and descending (D) orbits in non-forested, forested and total SAR-visible areas for (a) the Monte Mater, (b) the Val Canaria and (c) the Brienz areas.**

6.2 Monitoring Performance

460 Terrestrial data are used to characterise the ability of time series of displacements derived from C- and L-band multitemporal DInSAR products to follow the evolution of surface displacements at specific locations, where *in-situ* monitoring data is available. At Monte Mater, validation is possible due to 10 GNSS benchmarks maintained by ARPA Lombardia, with standard errors of 3–5 mm in horizontal and 6–10 mm in vertical components, and a Ku-band GB-DInSAR system

positioned 1.4–3.4 km from the slope, with a range resolution of 1.5 m and azimuth resolution of 5.8–11.5 m. To extract the PSI time series, we defined circular sampling areas with a 50 m radius centred on the GB-DInSAR/GNSS benchmarks. For comparison, considering the uncertainties related to different LOS (satellite versus terrestrial DInSAR) and the slope-parallel trends of GNSS measurements, all datasets were projected along the steepest slope (Notti et al., 2014). For sectors moving at less than 5 cm per year, L-band PSI provides displacement time series in close agreement to GNSS and GB-DInSAR, capturing multi-year accelerations that began in 2022 (Fig. 13a). C-band time series are more affected by sparse coverage and fail to reproduce these non-linear trends. At intermediate velocities of about 10 cm per year, SAOCOM-1 PSI continues to follow the slightly non-linear behaviour measured by GB-DInSAR, although intra-annual variations are not fully resolved (Fig. 13b). At higher velocities (i.e., ~20 cm per year), all satellite PSI products underestimate motion: C-band captures only about 25% of the true displacement, while L-band recovers around 50% and shows only a weak indication of non-linear acceleration (Fig. 13c). Tables of the theoretical velocity thresholds as discussed in Manconi (2021) are included in Appendix D. The thresholds indicate the maximum theoretical velocity along the LOS that can be detected by each satellite and thus express each sensor's upper velocity limits.

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At Brienz/Brinzauls, three continuous GNSS stations are located in the lower part of the slope, but only one point meets the predefined criteria for validation (PSI measurements within a 50 m buffer around the *in-situ* measurement). In 2018–2019, for LOS rates of motion of about 5 cm per year, the L-band time series matches GNSS displacements projected onto the SAR LOS (Fig 13d), confirming the reliability of the method under moderate velocities. After 2019, however, slope acceleration reached several metres per year, well beyond the measurable range of PSI. L-band results then strongly underestimate the real displacements and the fastest slope sectors yielded no PSI points at all. Here, the dense terrestrial network of GNSS, GB-DInSAR, and car-borne GLSAR provided the critical information, recording highly non-linear accelerations and mapping internal slope compartments with velocities of up to several metres per year.

485 The Rütan dei Sass rockfall in Val Canaria allows further quantitative assessment through annual Motorized Total Station (MTS) surveys, carried out since 2013. We have selected two MTS benchmarks: pt. 26 in the fast-moving central part of the landslide, where results from the SAOCOM-1 SBAS analysis are available (Fig. 8) and pt. 25 in the upper part of the landslide, which provides a high number of slow moving ALOS-2 PALSAR-2 and SAOCOM-1 PS data (Fig. 6). The satellite data is averaged within a radius of 50 metres around each MTS target. The displacements observed in the geodetic measurements are then projected onto the SAR LOS to enable an effective comparison. At pt. 26, the SAOCOM-1 SBAS displacement time series aligns very well with the MTS measurements in LOS between 2020 and 2023 (Figure 13e). Nevertheless, the strong acceleration between 2023 and 2024, although detected with the SAOCOM-1 SBAS analysis, is underestimated. Phase unwrapping errors in the interferograms spanning this time period are likely the reason of the

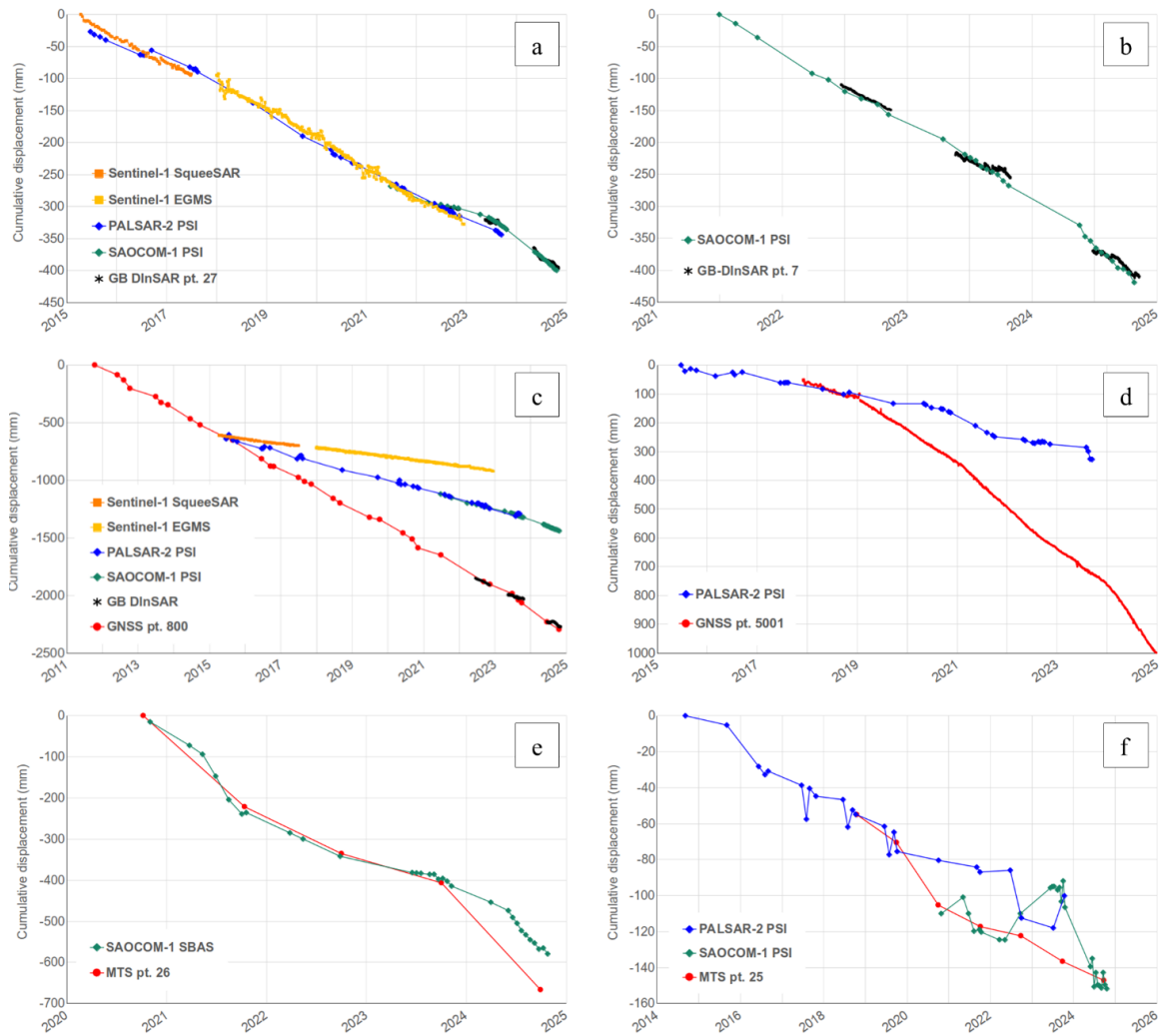
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discrepancy. At pt. 25 the slow displacement rates measured by MTS (15 mm/a between 2018 and 2024 and 10 mm/a
495 between 2020 and 2024) compare well with those observed by ALOS-2 PALSAR-2 and SAOCOM-1 (12 mm/a between
2014 and 2024 and 9 mm/a between 2020 and 2024, respectively). However, the time series of Figure 13f show more noise
and general deviations, e.g., in the period from October 2019 to October 2020, when a significant acceleration detected by
the MTS was not recorded by ALOS-2 PALSAR-2, and in 2022 and 2023, when SAOCOM-1 showed a misalignment
compared to the MTS data. Displacement rates of about 9 mm/a correspond to the expected detection accuracy threshold for
500 L-band (Ng et al., 2012), while the accuracy for single measurements is on the order of 15 mm (Nishiguchi et al., 2017).

6.3 Towards Multi-Frequency Integration

The results and validation thereof clearly illustrate the advantages and limitations of satellite and terrestrial L-band datasets
for DInSAR techniques in alpine environments. Each of the three study sites highlights important aspects of our findings and
provides insight into the monitoring and early warning capabilities of current and future L-band SAR missions.

505 The Monte Mater study site in Italy highlights the high potential of L-band to act as complementary data to the widely
available C-band Sentinel-1 datasets. While some information can be extracted from C-band PSI results due to the relatively
slow motion of the landslide and limited vegetation cover, the point coverage remains too scarce to clearly outline the full
extent of the moving area. L-band results from ALOS-2 PALSAR-2 and SAOCOM-1 strongly complement the C-band data
as their point density remains high even in areas of dense vegetation and fast movements. L-band PSI thus enables an
510 understanding of the full extent of the Monte Mater landslide and its compartments through identification of different
velocities within the moving area.



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Figure 13: Time series of displacements, extracted and averaged in circular sampling areas ($R=50\text{m}$) for different satellite C- and L-band multitemporal DInSAR datasets. a) Slow-moving, deep-seated rockslide at GB DInSAR pt. 27 at Monte Mater (rates $<5\text{ cm/a}$); b) shallow debris covering the Monte Mater DSGSD head at GB DInSAR pt. 7, moving at rates $\sim 10\text{ cm/a}$; c) shallow debris slide in the upper part of the Monte Mater DSGSD at GNSS pt. 800, moving at long-term rates $>20\text{ cm/a}$; d) GNSS pt. 5001 at Brienz village; e) MTS pt. 26 over the fast moving central part of the Rütan dei Sass rockfall in Val Canaria (long-term rates $\sim 20\text{ cm/a}$); f) MTS pt. 25 in the upper slow part of the Rütan dei Sass rockfall in Val Canaria (rates $\sim 1\text{ cm/a}$). At Monte Mater all datasets are projected along the steepest slope, for Brienz and Val Canaria the *in-situ* data are converted towards the satellite LOS. For Brienz values are shown in reverse direction for illustrative purposes.

The Val Canaria study site in Switzerland is a very good example of the obvious limitations of C-band to detect fast movements in vegetated areas and the clear need for lower frequency datasets with high temporal sampling rates. No information can be extracted from Sentinel-1 PSI results at Val Canaria, while L-band PSI provides a good solution and

525 indication of local displacement. The influence of a higher temporal sampling rate of available acquisitions becomes
apparent here through the different results obtained with ALOS-2 PALSAR-2 and SAOCOM-1. The higher temporal
sampling rate of SAOCOM-1 gives an overall higher point density compared to ALOS-2 PALSAR-2 and thus allows better
detection of displacement also in fast regions of slope movement. Despite this, gaps remain which are related to the limited
velocity range generally detectable with PSI. These can be filled with the use of differential interferograms and other
530 advanced DInSAR techniques such as SBAS. The high temporal sampling rate of SAOCOM-1 allows generation of a
displacement time series from SBAS results that clearly shows non-linear motion and seasonal variations in the landslide's
velocity. These findings suggest that with a sufficiently high temporal sampling, satellite-based L-band data can provide a
good understanding of inter-annual changes in displacement rates and may act as valuable information in monitoring
programmes.

535 The Monte Mater and Val Canaria landslides further confirm the potential of terrestrial L-band measurements as acquired
with the GLSAR instrument. The displacement maps obtained with the car-borne instrument are very similar to the satellite-
based results, even in Val Canaria despite difficult local conditions, such as a rough road along which the car is driven to
acquire measurements of the opposite slope. These findings support the high robustness of terrestrial measurements and
show the reciprocal validation potential of satellite- and ground-based data studies.

540 The Brienz study site in Switzerland strongly highlights the potential of multi-sensor and multi-approach studies to obtain a
good understanding of very fast-moving landslides with complex local kinematics. Again, the limitations of C-band and
advantages of L-band are illustrated by comparing the PSI point coverage of each frequency, where ALOS-2 PALSAR-2 and
SAOCOM-1 show distinctly higher point densities than Sentinel-1 and enable the identification of large areas of motion in
the surrounding region. The limitations of PSI regardless of the sensor wavelength, however, are also apparent, as the Brienz
545 landslide itself cannot be detected with this technique even at L-band. A multi-technique approach through use of differential
interferograms and other advanced DInSAR techniques is thus necessary to gain an understanding of the landslide
kinematics. In particular, the use of offset-tracking approaches could add valuable insight into the analysis of the Brienz
landslide (Manconi et al., 2024). Additionally, these findings show the current limitations of a purely satellite-based analysis,
as the fastest areas of the landslide in the upper parts of the slope, where velocities exceed several metres per year, cannot be
550 resolved with the given temporal sampling rate of approx. two weeks. These areas are coherently detected with terrestrial L-
band measurements acquired with the GLSAR, however, which are conducted within a time frame of one week and provide
valuable additional information on the upper parts of the landslide. The car-borne GLSAR clearly allows identification of the
landslide outline at these higher altitudes as well as of internal compartments with high velocities of several metres per year.
Satellite data available with sub-daily revisit times and at sub-metric spatial resolutions can also be a viable alternative to
555 manage slope instability operational scenarios (Manconi et al., 2024).

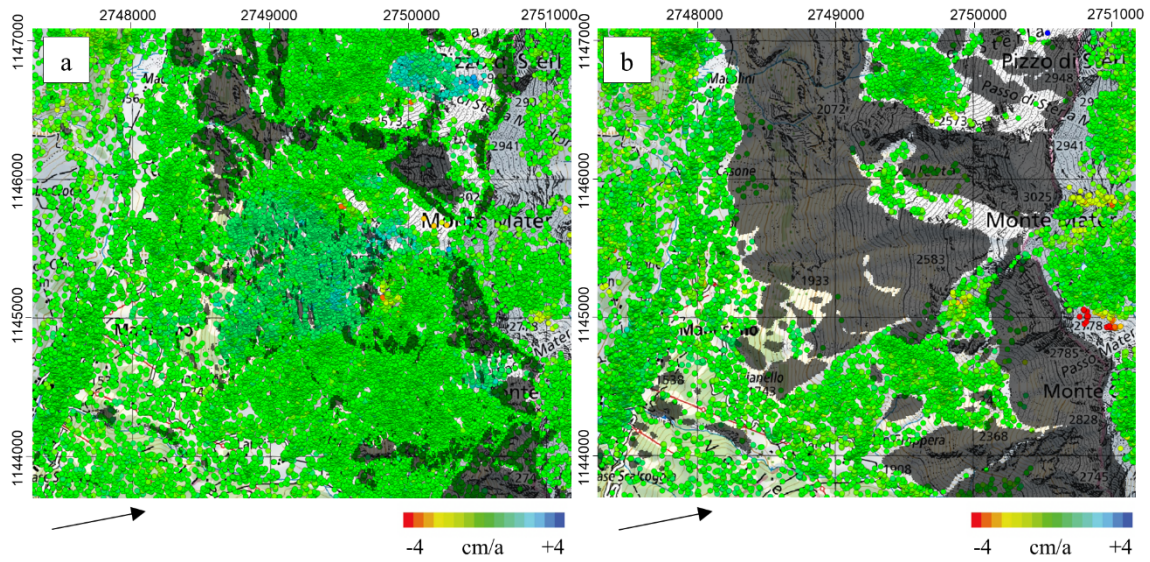
7 Conclusion and Outlook

Our comparative analyses at Monte Mater, Val Canaria, and Brienz/Brinzauls demonstrate the added value of L-band interferometry for both the detection and monitoring of slope instabilities in alpine environments. The PS densities achieved at L-band are systematically higher than at C-band, by a factor of five in forested terrain to a factor of more than twenty in favourable conditions. This improvement directly translates into more complete spatial coverage and the ability to delineate kinematic compartments that remain unresolved in C-band datasets. The case of Val Canaria is particularly relevant. While Sentinel-1 offers negligible coverage in forested areas, ALOS-2 and SAOCOM-1 provide dense scatterer distributions across both vegetated and non-vegetated slopes, enabling detection of slope sectors that had previously gone unrecognised. At Brienz, PSI at both frequencies fails to capture the core of the landslide once velocities exceed measurable thresholds. These results confirm and extend earlier observations of C-band limitations in alpine terrain, providing quantitative evidence that L-band coherence and scatterer density can systematically overcome vegetation-related decorrelation. The monitoring capabilities of L-band are more dependent on velocity and acquisition frequency. For slow slope movements below ~5 cm/year, L-band PSI time series are consistent with GNSS and GB-DInSAR, including the ability to resolve multi-year accelerations. SAOCOM-1, with its higher temporal sampling, is able to track displacements of ~10 cm/year, although intra-annual variability observed by GB-DInSAR is not fully captured. At higher velocities, all satellite PSI products underestimated motion: C-band recovered only about 25% and L-band about 50% of the true displacement at Monte Mater, with similar patterns observed in Brienz. These findings highlight a fundamental limitation of interferometric monitoring in rapidly deforming environments, where terrestrial measurements remain essential for reliable early warning.

Together, our findings indicate that C- and L-band missions should be viewed as complementary rather than competing assets. Sentinel-1 provides broad regional coverage for baseline detection of slow and steady instabilities, while L-band sensors are indispensable for resolving displacements in vegetated areas and at moderate velocities. Nevertheless, the relatively low revisit frequency of current L-band missions continues to limit their usefulness for detecting rapid or short-term displacement phenomena. The launch of NISAR in 2025 and ROSE-L later this decade will provide, for the first time, free and regular L-band coverage at the global scale. Our results suggest that these missions will substantially enhance the capacity for regional landslide mapping and site-specific monitoring in the Alps. By systematically exploiting L-band's improved coherence and velocity detection range, it will be possible to achieve more accurate assessments of slope activity, long-term kinematic trends, and the influence of changing climatic drivers. However, the intrinsic limitations of interferometry at high displacement rates mean that satellite observations must be embedded within multi-sensor monitoring frameworks to ensure effective early warning and risk reduction.

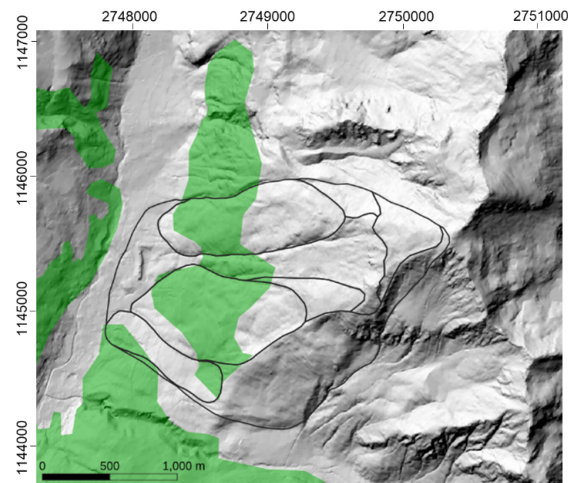
Table A1. List of L-band satellite acquisitions used in the PSI analysis of Monte Mater. ALOS-2 PALSAR-2 ScanSAR data are shown in *italics*. Additionally, the mean SAR look vector incidence angle (*lv_theta*) and mean SAR look vector orientation (*azimuth*) angle (*lv_phi*) are included per sensor.

ALOS-2 PALSAR-2 A199-F2-7	ALOS-2 PALSAR-2 D095-F2-5	SAOCOM-1 Asc S3 T214	SAOCOM-1 Desc S4 T112
<i>lv_theta: 41.8°</i>	<i>lv_theta: 33.5°</i>	<i>lv_theta: 31.8°</i>	<i>lv_theta: 36.2°</i>
<i>lv_phi: -170.5°</i>	<i>lv_phi: -10.7°</i>	<i>lv_phi: -169.1°</i>	<i>lv_phi: -10.4°</i>
20140906	<i>20150625</i>	20210806	20210702
20150905	<i>20150723</i>	20210923	20210819
20160611	<i>20150903</i>	20220505	20211022
<i>20160806</i>	<i>20151015</i>	20220622	20220331
20160903	<i>20160623</i>	20220825	20220518
20170624	20160707	20220910	20220705
<i>20170805</i>	<i>20160721</i>	20221012	20220822
20170902	20160915	20221113	20220907
20180623	<i>20170622</i>	20230321	20221009
20180804	<i>20170720</i>	20230508	20221110
<i>20180915</i>	20170803	20230524	20230419
20181013	<i>20170817</i>	20230617	20230622
20190622	<i>20180927</i>	20230703	20230708
20190803	<i>20190926</i>	20230719	20230724
<i>20190914</i>	20200423	20230804	20230809
20191012	20200507	20230820	20230825
20201010	20200521	20230905	20230910
<i>20210911</i>	20200702	20230921	20230926
20211009	20200910	20231007	20231012
20220702	20200924	20231023	20231028
20221008	20201105	20231108	20240523
<i>20230715</i>	20201119	20231210	20240608
20231021	20210812	20231226	20240624
	20210923	20240416	20240710
	20211007	20240502	20240726
	20220519	20240518	20240811
	20220714	20240619	20240827
	20220728	20240705	20240912
	20220825	20240721	20240928
	20220908	20240806	20241014
	<i>20220922</i>	20240822	20241030
	20221006	20240907	
	20221117	20240923	
	<i>20230810</i>	20241110	
	<i>20230824</i>	20241126	
	20230907	20241212	
	20230921	20241228	



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Figure A2. Results of PSI analysis at Monte Mater with ascending ALOS-2 PALSAR-2 for the time period 2014–2023 (a) and SAOCOM-1 for the time period 2021–2024 (b). The arrows indicate the satellite LOS direction. Layover/shadow is shown in black. Image background is the Swiss National Map.



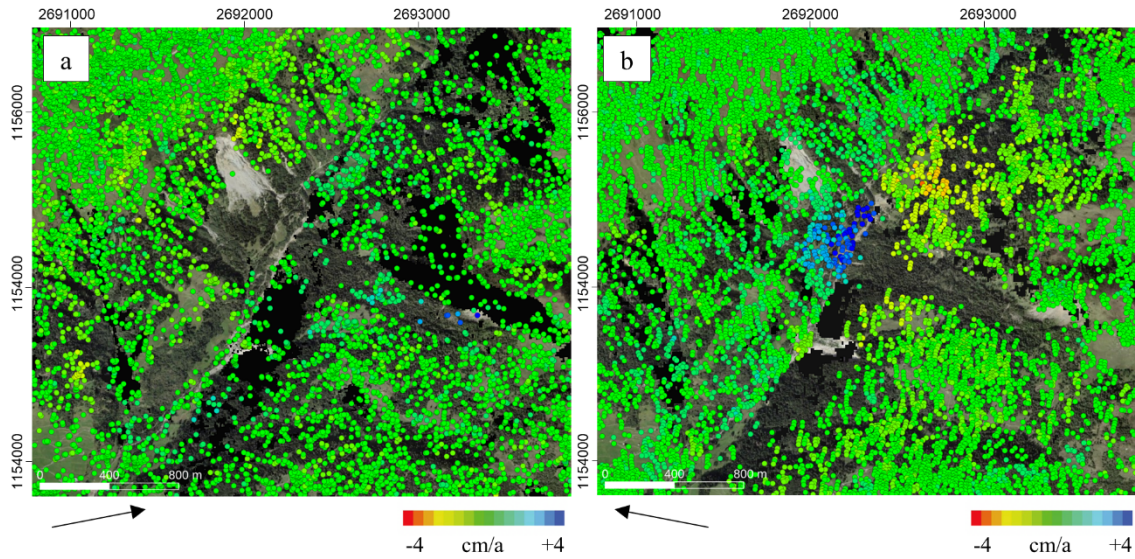
595

Figure A3. Forest cover (green) from the CORINE Forest Land (CLC, <https://land.copernicus.eu/en/products/corine-land-cover/>, last accessed 31 March 2026) and outlines of the landslide complex (black; Crippa et al., 2020) at Monte Mater. Image background is the DEM shaded relief.

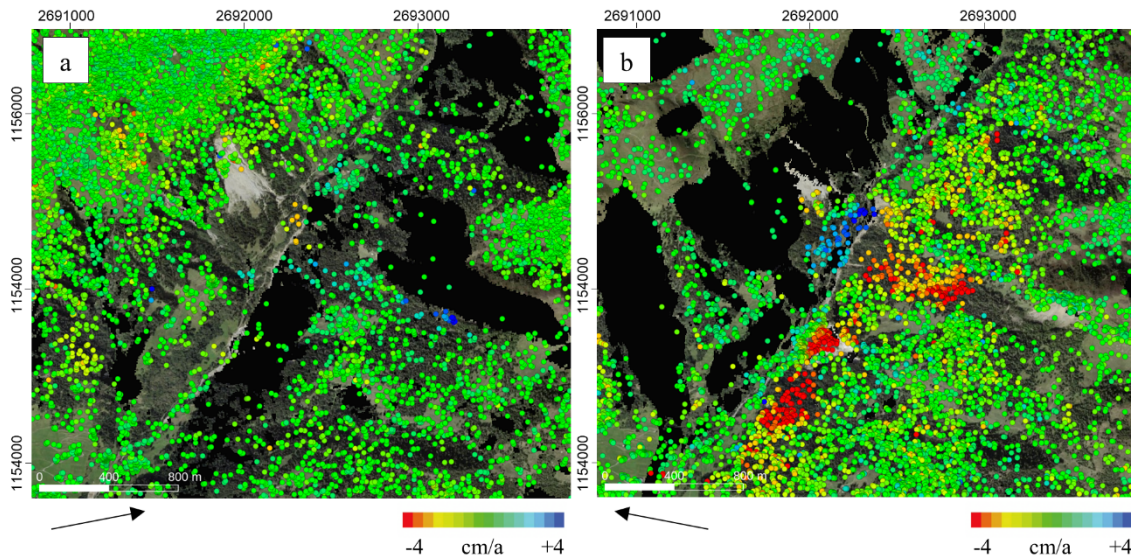
Appendix B: Val Canaria, Ticino, Switzerland

600 **Table B1.** List of L-band satellite acquisitions used in the PSI analysis of Val Canaria. ALOS-2 PALSAR-2 ScanSAR data are shown in italics. Additionally, the mean SAR look vector incidence angle (lv_theta) and mean SAR look vector orientation (azimuth) angle (lv_phi) are included per sensor.

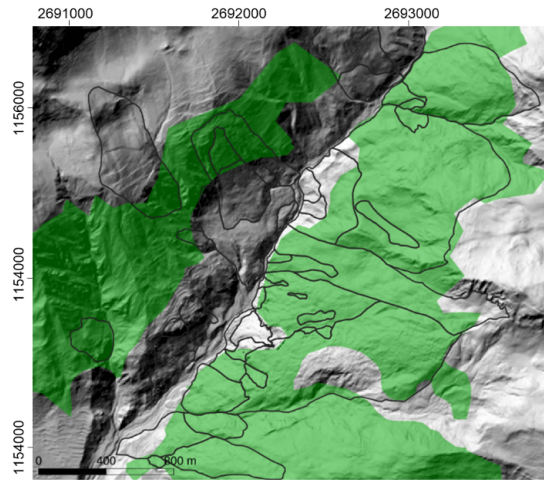
ALOS-2 PALSAR-2 A199-F2-7	ALOS-2 PALSAR-2 D095-F2-6	SAOCOM-1 Asc S4 T215	SAOCOM-1 Desc S3 T113
lv_theta: 38.9°	lv_theta: 37.6°	lv_theta: 36.3°	lv_theta: 32.0°
lv_phi: -170.0°	lv_phi: -10.2°	lv_phi: -169.6°	lv_phi: -10.9°
20140906	<i>20150625</i>	20201102	20220825
20150905	<i>20150723</i>	20210513	20220910
20160611	<i>20150903</i>	20210630	20221012
<i>20160806</i>	<i>20151015</i>	20210817	20221113
20160903	<i>20160623</i>	20211004	20221129
20170624	<i>20160721</i>	20211020	20230116
<i>20170805</i>	20170525	20220329	20230201
20170902	<i>20170622</i>	20220516	20230422
20171028	<i>20170720</i>	20220929	20230703
20180623	<i>20170817</i>	20230620	20230804
20180804	20180524	20230706	20230905
<i>20180915</i>	20180607	20230722	20231007
20181013	<i>20180927</i>	20230823	20231108
20190622	20190523	20230908	20240212
20190803	20190606	20230924	20240315
<i>20190914</i>	<i>20190926</i>	20231010	20240416
20191012	20200604	20231026	20240518
20201010	20210603	20240606	20240619
<i>20210911</i>	20220602	20240622	20240721
20211009	20230601	20240708	20240822
20220702	<i>20230810</i>	20240724	20240923
20221008	<i>20230824</i>	20240809	
<i>20230715</i>		20240825	
20231021		20240910	
		20240926	
		20241012	
		20241028	



605 **Figure B2. Results of PSI analysis at Val Canaria with ascending ALOS-2 PALSAR-2 for the time period 2014–2023 (a) and descending ALOS-2 PALSAR-2 for the time period 2015–2023 (b). The arrows indicate the satellite LOS direction. Layover/shadow is shown in black. Image background is SWISSIMAGE.**



610 **Figure B3. Results of PSI analysis at Monte Mater with ascending SAOCOM-1 for the time period 2020–2024 (a) and descending SAOCOM-1 for the time period 2022–2024 (b). The arrows indicate the satellite LOS direction. Layover/shadow is shown in black. Image background is SWISSIMAGE.**



615 **Figure B4.** Forest cover (green) from the CORINE Land Cover (CLC, <https://land.copernicus.eu/en/products/corine-land-cover/>, last accessed 31 March 2026) and inventoried slope instabilities (black; De Pedrini, personal communication, 2025) in the Val Canaria. Image background is the DEM shaded relief.

Appendix C: Brienz/Brinzauls, Grisons, Switzerland

Table C1. List of L-band satellite acquisitions used in the PSI analysis of Brienz/Brinzauls. ALOS-2 PALSAR-2 ScanSAR data are shown in *italics*. Additionally, the mean SAR look vector incidence angle (lv_theta) and mean SAR look vector orientation (azimuth) angle (lv_phi) are included per sensor.

ALOS-2 PALSAR-2 A199-F3-8	ALOS-2 PALSAR-2 D095-F2-5	SAOCOM-1 Asc S3 T214	SAOCOM-1 Desc S4 T112
lv_theta: 43.5°	lv_theta: 32.8°	lv_theta: 33.4°	lv_theta: 35.3°
lv_phi: -170.7°	lv_phi: -10.8°	lv_phi: -169.3°	lv_phi: -10.5°
<i>20160806</i>	<i>20150625</i>	20210619	20220822
<i>20170805</i>	<i>20150723</i>	20210806	20220907
<i>20180915</i>	<i>20150903</i>	20210923	20221009
<i>20190914</i>	<i>20151015</i>	20220318	20221110
20200606	20160303	20220505	20230113
20200620	20160707	20220622	20230129
20200801	<i>20160721</i>	20220825	20230214
20200829	20160915	20220910	20230419
20200912	<i>20170622</i>	20221012	20230708
20201024	<i>20170720</i>	20221113	20230809
20201107	20170803	20221215	20230910
20210703	<i>20170817</i>	20230116	20231012
20210717	20180426	20230321	20240217
20210814	<i>20180927</i>	20230508	20240320
20210828	<i>20181108</i>	20230524	20240523
<i>20210911</i>	<i>20190926</i>	20230617	20240624
20210925	20200507	20230703	20240726
20220604	20200521	20230719	20240827
20220618	20200702	20230804	20240928
20220716	20200910	20230820	20241030
20220730	20200924	20230905	20241201

20220827	20201105	20230921	
20220910	20201119	20231007	
20230617	20210520	20231023	
20230701	20210812	20231108	
20230715	20210923	20240212	
20230909	20211007	20240315	
20230923	20220505	20240416	
	20220519	20240502	
	20220714	20240518	
	20220728	20240619	
	20220825	20240705	
	20220908	20240721	
	20220922	20240806	
	20221006	20240822	
	20221117	20240907	
	20230810	20240923	
	20230824	20241110	
	20230907	20241126	
	20230921	20241212	
		20241228	

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Table C2. Ascending SAOCOM-1 and descending ALOS-2 PALSAR-2 differential interferograms used for stacking at the Brienz study site.

SAOCOM-1 Asc S3 T214	ALOS-2 PALSAR-2 D095-F2-5
20240315 20240331	20240404 20240418
20240331 20240416	20240418 20240516
20240416 20240502	20240627 20240711
20240502 20240518	20240711 20240808
20240619 20240705	20240808 20240822
20240705 20240721	20240822 20240919
20240907 20240923	20240919 20241003

Table C3. GLSAR measurements acquired at the Brienz study site.

GLSAR Measurements
20200120, 20200124, 20200203
20230626, 20230627, 20230628, 20230629
20231016, 20231017, 20231023, 20231030
20240820, 20240822

625

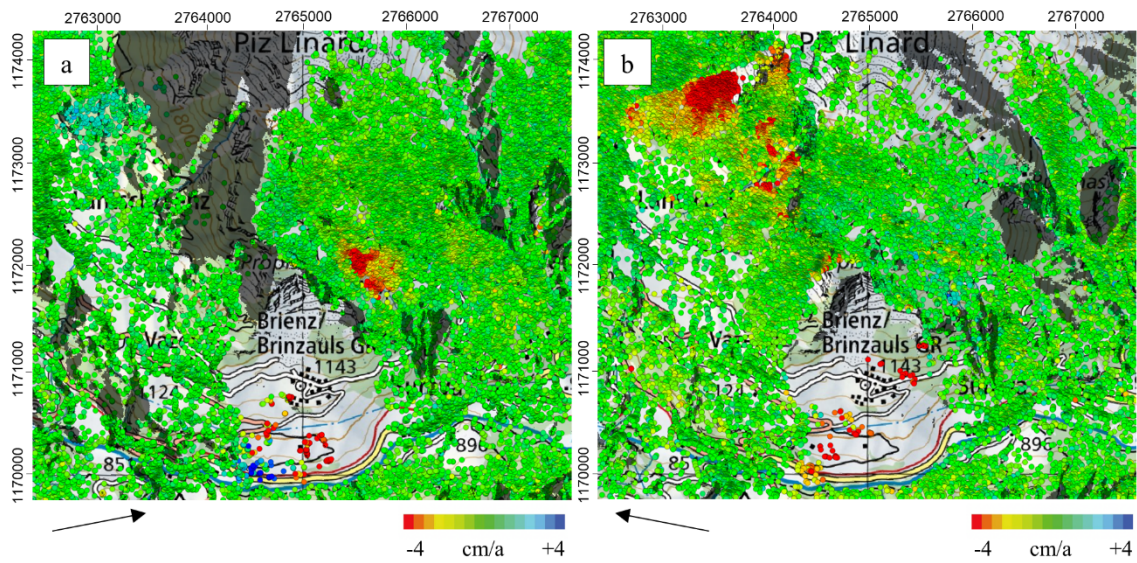


Figure C4. Results of PSI analysis at Brienz with ascending SAOCOM-1 for the time period 2021–2024 (a) and descending SAOCOM-1 for the time period 2022–2024 (b). The arrows indicate the satellite LOS direction. Layover/shadow is shown in black. Image background is the Swiss National Map.

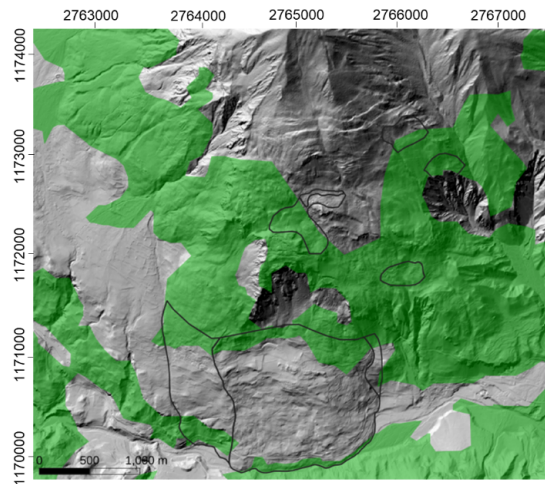


Figure C5. Forest cover (green) from the CORINE Land Cover (CLC, <https://land.copernicus.eu/en/products/corine-land-cover/>, last accessed 31 March 2026) and outlines of landslide hazards (black; Geoportail Canton Grisons, <https://map.geo.gr.ch>, last accessed 27 February 2026) at Brienz/Brinzauls. Image background is the DEM shaded relief.

Appendix D: Temporal phase aliasing thresholds per sensor

635 **Table D4: Theoretical velocity thresholds for standard DInSAR per sensor, calculated based on nominal wavelengths and revisit times (Manconi, 2021). These temporal phase aliasing thresholds refer to the case where the displacement of the target under investigation exceeds the threshold value of $\lambda/4$ between two consecutive satellite acquisitions.**

Sensor	Frequency [Hz]	Wavelength [m]	Interval [days]	Threshold [m/year]
ALOS-2 PALSAR-2	1.24E+09	0.243	14	1.582
SAOCOM-1	1.28E+09	0.235	16	1.342
GLSAR	1.30E+09	0.231	7	3.010
GLSAR	1.30E+09	0.231	82	0.257
GLSAR	1.30E+09	0.231	239	0.088

640 **Table D5: Theoretical velocity thresholds considering 28 days ALOS-2 PALSAR-2 used in the Brienz DInSAR stacking analysis (cf. Fig. 10), 270 days SAOCOM-1 used in the Val Canaria SBAS analysis (cf. Fig. 8) and 365 days intervals used for PSI analyses, calculated based on nominal wavelengths and revisit times (Manconi, 2021). These temporal phase aliasing thresholds refer to the case where the displacement of the target under investigation exceeds the threshold value of $\lambda/4$ during the acquisition time intervals of the data used in each processing method.**

Sensor	Frequency [Hz]	Wavelength [m]	Interval [days]	Threshold [m/year]
Sentinel-1	5.41E+09	0.056	365	0.014
ALOS-2 PALSAR-2	1.24E+09	0.243	28	0.791
ALOS-2 PALSAR-2	1.24E+09	0.243	365	0.061
SAOCOM-1	1.28E+09	0.235	270	0.080
SAOCOM-1	1.28E+09	0.235	365	0.059

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Contributions

T.S. and A.M. conceived the study. T.S., N.J. and P.B. performed the satellite DInSAR processing. O.F. and R.C. performed the terrestrial measurements and DInSAR processing. F.A., A.D.P., C.A. and A.M. designed the validation methodology and interpreted the results. T.S. and N.J. wrote the paper. F.A., A.D.P., O.F., P.B., R.C., C.A. and A.M. reviewed the manuscript.

655 **Declaration of Competing Interest**

The authors declare that they have no conflict of interest.

Data Sources

Digital elevation models (DEM) at high spatial resolution are available for all three study sites through the regional or national authorities. For Monte Mater in Italy, the DTM 5X5, 2015 edition, is made available through Lombardia region's
660 online geoportal Geoportale della Lombardia (<https://www.geoportale.regione.lombardia.it/>, last accessed 17 April 2024). For the two Swiss sites at Val Canaria in Canton Ticino and Brienz in Canton Grisons, the 2m and 0.5m resolution SwissAlti3D DEM is publicly available through swisstopo of the Federal Office of Topography (swissALTI3D: <https://www.swisstopo.admin.ch/de/hoehenmodell-swissalti3d/>, last accessed 17 April 2024).

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