https://doi.org/10.5194/egusphere-2025-5106
Preprint. Discussion started: 24 November 2025

The multi-decadal hazard cascade of a tropical mountain wildfire

William Veness[1], Martha Day[1], Anthony C. Ross[1], Yazidhi Bamutaze[2], Jiayuan Han[1], Douglas Mulangwa[3,4], Andrew Mwesigwa[4], Emmanuel Ntale[5], Callist Tindimugaya[4], Brian Guma[4], Elisabeth Stephens[3,6], Wouter Buytaert[1]

[1]Department of Civil and Environmental Engineering, Imperial College London, London, SW7 2BB, UK

[2]Department of Geography, Geo-Informatics and Climatic Sciences, Makerere University, Kampala, Uganda

[3]Department of Meteorology, University of Reading, Reading, RG6 6BB, UK

[4]Ministry of Water and Environment, Kampala, Uganda

[5]Uganda Red Cross Society, Kampala, Uganda

[6]Red Cross Red Crescent Climate Centre, The Hague, The Netherlands

Correspondence to: William Veness (williamaveness@gmail.com)

Abstract. Climate change is driving wildfires to higher elevations, yet the hazard cascades that follow the burning of pristine tropical mountain ecosystems remain largely unexplored. Here, we analyse the long-term cascade following a February 2012 wildfire that burned 31 km² of forest and wetland in Uganda's Rwenzori Mountains National Park. Combining remote sensing, humanitarian records, field surveys, and interviews, we document ten major floods since 2012, including two debris floods that required large-scale humanitarian responses. Post-fire increases in erosion and mass movement have widened the River Nyamwamba sevenfold since 2012, breaching copper-cobalt mine tailings and mobilising an estimated 744,000 tonnes of waste into the river. Slow vegetation recovery at high altitudes and positive feedbacks between hazards have prolonged this high-risk state, underscoring the susceptibility of tropical mountain ecosystems to long-term post-wildfire cascades. More monitoring and research are required to characterise key hazard interactions after tropical mountain fires, which can guide entry points for management seeking to mitigate and impede future cascades.

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1 Introduction

Climate and land-use changes are driving more frequent and intense wildfires across many tropical ecosystems worldwide (Ometto et al., 2022; UNEP, 2022; Wimberly et al., 2024; Obando-Cabrera et al., 2025). In tropical mountains, fires are burning at higher elevations (Xiao et al., 2022), which is exposing mature forests and wetlands that are not adapted to burning regimes. Tropical mountain forests cover 1.8 million km² globally (FAO & UNEP, 2020). They provide the headwaters of major river systems such as the Nile, Amazon and Mekong, and sustain the livelihoods of over 336 million people (Encalada et al., 2019).

Tropical mountains host multiple hazards, making them susceptible to multi-hazard cascades (Arango-Carmona et al., 2025). Intense convectional rainfall drives flash floods; high temperatures at lower elevations cause droughts, heatwaves and wildfires (Ometto et al., 2022); and they are tectonically active (Sandwell et al., 2005). Their steep gradients, deeply chemically weathered soils, and unconsolidated glacial and fluvial deposits also favour landslides, debris flows and erosion (Arango-Carmona et al., 2025). Multi-hazard cascades occur when two or more of these hazards interact through multiple relationships characterised as *triggering*, *probability increasing*, or *catalysing/impeding* (Gill and Malamud, 2016).

1.1 Post-Wildfire Hazard Cascades

Despite their increasing risk, wildfire hazard cascades in tropical montane regions remain poorly understood. Most existing research comes from temperate systems, where wildfires are known to amplify floods, accelerate erosion, and increase the probability of landslides and debris flows by removing vegetation, altering soil properties and increasing surface runoff (Belongia et al., 2023; Boyer et al., 2022; DeBano, 2000; Doerr et al., 2000; Guerriero et al., 2025; Jordan, 2016; Kemter et al., 2021; Rengers et al., 2020; Swain et al., 2025; Vahedifard et al., 2024).

However, there are additional factors in tropical mountains that introduce greater risk and complexity (Moazeni & Cerdà, 2024; Robinne et al., 2021). First, the fires impact upon an already diverse multi-hazard landscape with many existing hazard interactions (Arango-Carmona et al., 2025; Ometto et al., 2022; Sandwell et al., 2005). Second, many higher-altitude ecosystems within tropical mountains have no history of wildfire, such that mature climax vegetation and wetlands are burned with unpredictable consequences for hydrological processes and ecosystem services (Marengo et al., 2021; Pivello et al., 2021; UNEP, 2022). Third, a lack of wildfire history means vulnerable populations without lived experience are exposed to new hazards (McCaffrey, 2004; Paton, 2003). Lastly, vegetation recovery at high altitudes is slow due to cold conditions, a thinner atmosphere, and the presence of vegetation that is not adapted to fire cycles, causing prolonged impacts (Kappelle et al., 1996; Oliveras et al., 2014; Salinas et al., 2021). Given these differences, there is a need to better understand the long-term cascade of tropical montane wildfires at the process level. This is especially true for multi-hazard risk management, as identifying where hazards interact effectively highlights where those interactions can be proactively impeded (AghaKouchak et al., 2018, Aghakouchak et al., 2020; Vahedifard et al., 2024).

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1.2 Rwenzori Mountains National Park 2012 Wildfire

The February 2012 wildfire in Uganda's Rwenzori Mountains National Park burned 31 km² of pristine, uninhabited tropical mountain forest and wetlands (Fig. 1) during a brief meteorological drought measuring -3.5 in a 1-month Standardised Precipitation Index (Appendix A). The fire was followed by unprecedented debris flooding in May 2013 that displaced more than 25,000 people, caused 13 deaths and over USD $4 million in damages (Delforge et al., 2025). Local rainfall records suggested only a 6.6-year return interval, indicating that post-fire landscape changes drove the disaster (Jacobs et al., 2016). More than a decade later, the Nyamwamba catchment continues to experience flooding, debris flows, mass movements, erosion and water pollution at elevated intensity. Because the wildfire occurred in a protected area with no burn history and little subsequent intervention (Norville, 2024), it provides an unparalleled case for this study to characterise the long-term multi-hazard cascade of a tropical mountain wildfire.

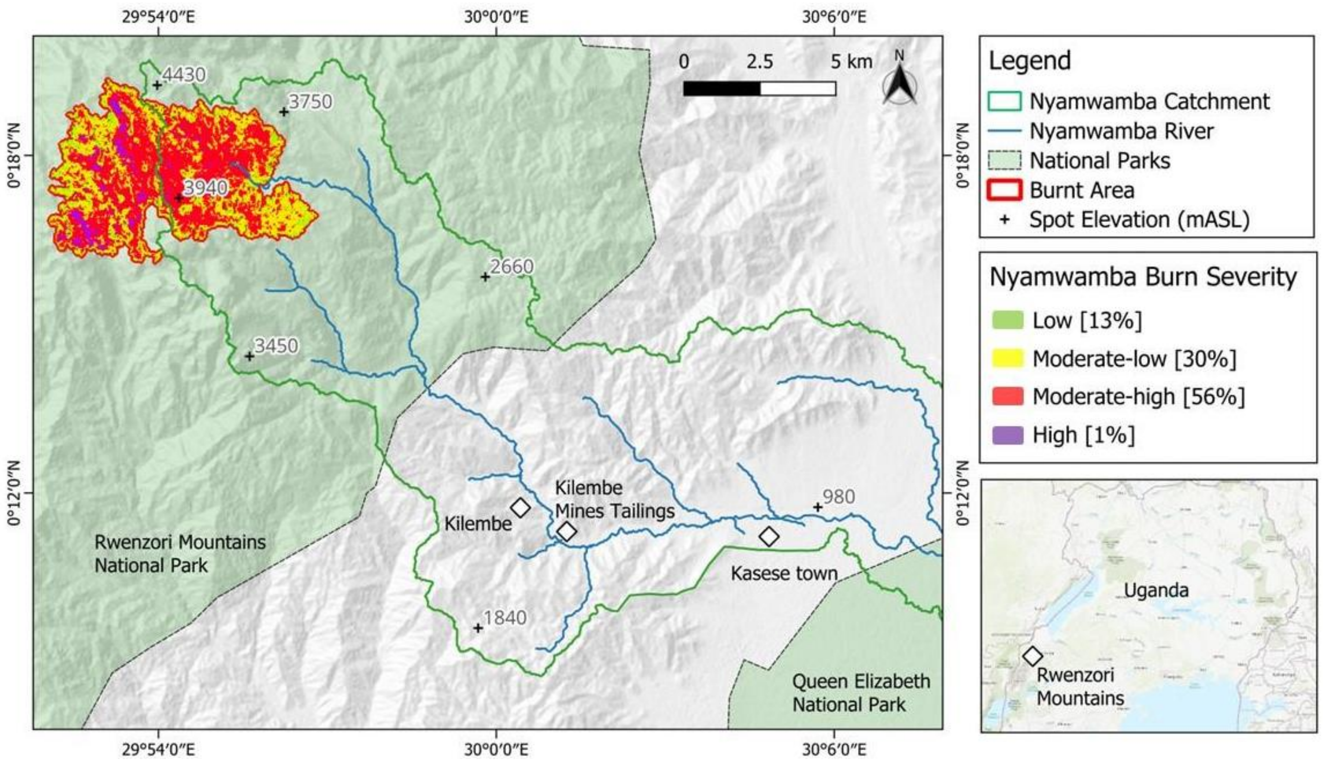

Figure 1: The River Nyamwamba catchment and the delineated wildfire burn area within the Rwenzori Mountains, Uganda. Differenced Normalised Burn Ratio (dNBR) between pre- and post-fire Landsat-7 images were used to delineate the extent and burn severity of the February 2012 wildfire. Severity is classified according to the United States Geological Survey's guide (Key & Benson, 2006).

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2 Methods

We adopted a mixed methods approach to evidence changes in multi-hazard processes and risk, combining remote sensing, humanitarian data, field observations and key informant interviews. Cross-validation across methods enabled an abductive approach (Saunders et al., 2016), where emerging insights, such as interview reports of erosion, informed subsequent lines of data collection and analysis.

2.1 Remote Sensing and GIS

2.1.1 Data Acquisition and Pre-Processing

Annual Landsat-7 (2006 – 2012) and Landsat-8 (2013-2024) Level-2 surface reflectance images were downloaded from the United States Geological Survey (USGS) earth explorer and gap corrected, cropped and cloud masked for analysis (Congedo, 2021). For each year, the earliest post-January 1 image with <10% cloud cover was selected. High-resolution Google Earth Pro imagery was used to measure river width, while Maxar mosaics visualised mine tailings erosion (Maxar Technologies, 2025a; Maxar Technologies, 2025b).

2.1.2 Burn Severity Classification

Burned area was delineated using the Normalised Burn Ratio (NBR), which combines Landsat 7 near-infrared (Band 4) and shortwave-infrared (Band 7) reflectance (Key & Benson, 2006). The difference between pre- (9 January 2012) and post-fire (28 March 2012) NBR values (dNBR) provided a relative severity index following USGS protocols (Key & Benson, 2006).

2.1.3 River Erosion Analysis

Supervised minimum-distance land-cover classifications were applied to annual Landsat images from 2006 – 2024, using fixed ground control points for five classes: eroded river channel, tailings, oxidised iron, vegetation, and agriculture (Congedo, 2021). Each image was clipped to the Nyamwamba channel, and classified areas were validated against Google Earth area estimates with a relative error of 3.84%. Cumulative eroded area was plotted over time, with classification maps from 2006 and 2021 shown for comparison. River width was delineated in 2010, 2014, 2018, and 2021, at 1 km intervals along 20 km of channel between Kilembe town and Lake George.

2.1.4 Tailings Erosion

Erosion of the Kilembe Mines tailings was assessed using Maxar mosaics from March 2006 and April 2023, with the 33,000 m² eroded footprint delineated manually. Field measurements in July 2024 using a laser rangefinder provided site dimensions, from which eroded volumes were calculated (see Appendix B).

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2.2 Humanitarian Data Analysis

Historic flood events in the Nyamwamba catchment since 2000 were compiled from multiple open sources: the Emergency Events Database (Delforge et al., 2025), the Sendai DesInventar database (DesInventar, 2025), grey literature in ReliefWeb, and a systematic keyword search ("Kasese" OR "Kilembe" AND "flood") across Google, Google Scholar, and Google News (Google News, 2025). While recent years benefit from expanded monitoring and reporting, the inclusion of diverse sources provided confidence that all major flood events since 2000 were captured by the search.

2.3 Interviews

We conducted twelve in-person semi-structured interviews during field visits in 2023 and 2024, following ethical clearance. Participants were identified through project partners in Kasese District, with snowball sampling to access other stakeholder groups. They included 2 representatives from the Ministry of Water [M – code used to reference in the results], 2 local government officials [G], 1 wildlife authority employee [W], 1 non-governmental organisation worker [N], 2 local industry workers [I], 1 farmer [F], and 3 community residents [R].

Interviews followed a lightly structured topic guide covering hazard processes and causes, changing risk, existing management, and potential alternatives, while remaining flexible to emergent themes (Creswell, 2009; Galletta, 2013; Mojtahed et al., 2014). A full guide is provided in Appendix C. Interviews lasted 30 – 120 minutes, were audio-recorded, transcribed, and coded inductively over two rounds of review, with related codes grouped into interpretive themes (Patton, 2014; Saldana, 2021). While themes are not presented directly, this analysis informed interpretation of hazard processes, impacts, and management options.

2.4 Photographs

Historic photographs of the vegetation pre- and immediately post-wildfire were taken by project partners with permission for research use. Photographs in Appendices E – J were taken by the study authors during a July 2024 site visit.

2.5 Cascade Visualisation

Processes identified through the above methods were integrated into a conceptual diagram of the wildfire's multi-hazard cascade (Patton, 2014), following Gill & Malamud's (2016) framework for hazard interaction types. Evidence underlying each connection is documented throughout the Results and summarised in Table D1 (Appendix D).

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3 Results

We present the multi-hazard cascade caused by the 2012 Rwenzori National Park wildfire (Fig. 2). The following sections describe each of the hazards involved and the interactions they drive, based on evidence from our mixed methods. Results are structured by hazard type: wildfire (Sect. 3.1), flooding (3.2), landslides (3.3), erosion (3.4), and pollution (3.5). Identified interactions highlight opportunities where management interventions can impede the cascade, for which we discuss practical solutions at the local and global scales in Sect. 4 (Discussion).

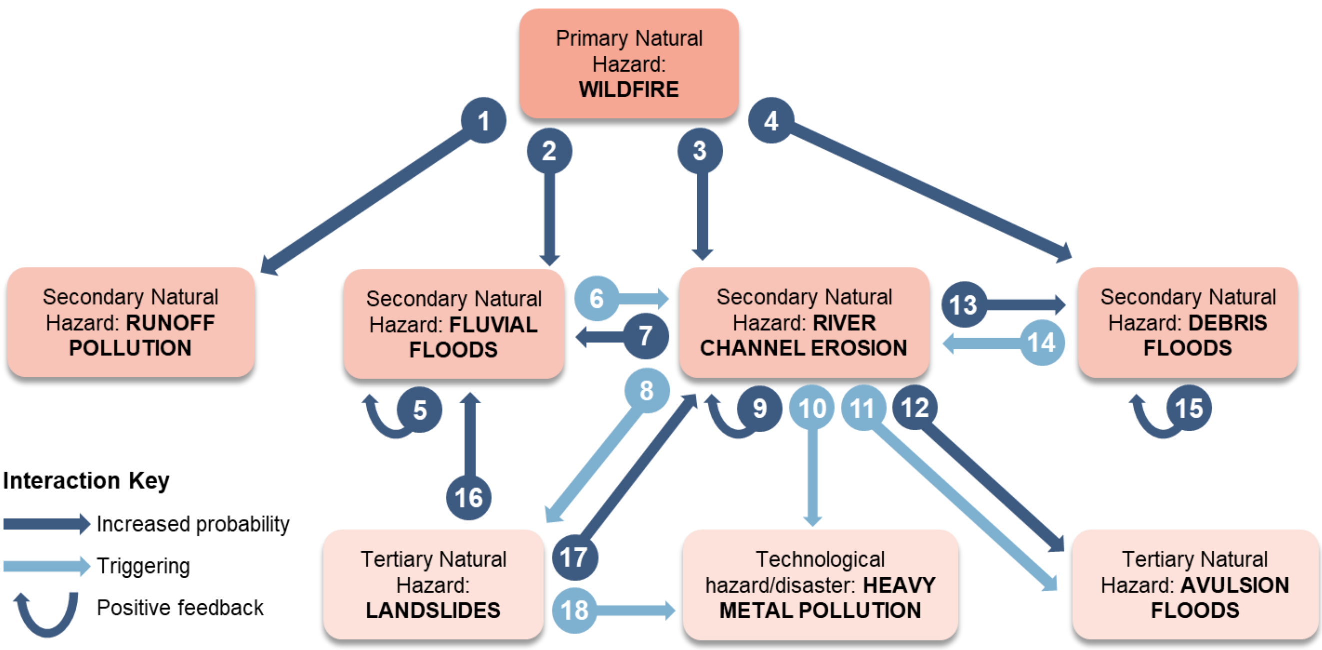

Figure 2: Conceptual model of the multi-hazard cascade following the Rwenzori National Park wildfire in February 2012. Interactions between hazards are classified as being (i) triggering, (ii) probability increasing or (iii) catalysing/impeding, following Gill and Malamud's (2016) framework. As there are numerous catalysing/impeding relationships in this context, we omit these from the visualisation for simplicity but describe key examples in the analysis text. Table 1 describes each of the interactions shown.

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Table 1: Description of the hazard cascade interactions in Fig. 2. The study evidence for each interactions is explained in the text and summarised in Table D1 (Appendix D).

Driving Hazard		Description of Interaction	Affected Hazard
Wildfire	1	Wildfire generated ash & exposed soils to surface runoff	Runoff Pollution
	2	Burning increased runoff & river discharge, causing higher peak flows	Fluvial Floods
	3	The higher peak river discharge has increased the river's erosive power	River Erosion
	4	The higher peak river discharge has increased its transport competence	Debris Floods
Fluvial Floods	5	Each flood damages natural banks & flood defences	Fluvial Floods
	6	Higher flow velocities & turbulence during floods increase erosion	River Channel Erosion
River Channel Erosion	7	Eroded material fills the channel, reducing its discharge capacity	Fluvial Floods
	8	Lateral erosion undercuts & destabilises hillslopes	Landslides
	9	Lateral erosion exposes bare river cliffs to further erosion	River Channel Erosion
	10	Direct erosion inputs Co-Cu Kilembe Mines solid tailings into the river	Heavy Metal Pollution
	11	Higher erosion rates have increased channel-switching events	Avulsion Floods
	12	Eroded sediment deposits in channel bars, diverting flow to banks	Avulsion Floods
	13	Erosion generates additional sediment for debris flow formation	Debris Floods
Debris Floods	14	Debris floods damage natural banks & flood defences	Debris Floods
	15	Debris flows have a high erosive power	River Channel Erosion
Landslides	16	Landslide talus fills the channel, reducing its discharge capacity	Fluvial Floods
	17	Landslides increase sediment loads, increasing erosion by abrasion	River Channel Erosion
	18	Rotational slumping of tailings inputs waste to the river channel	Heavy Metal Pollution

3.1 Wildfire

Remote sensing evidence shows a 30.75 km^2 burn area for the February 2012 wildfire (Fig. 1), with 87% of the area burned to a moderate or high severity. The fire occurred during a meteorological drought, with <0.2 mm of precipitation in the 4-weeks preceding the fire (Jacobs et al., 2016) and a one-month Standardised Precipitation Index measuring -3.5 for January 2012 (Appendix A). The trigger of the fire is still unknown by the water and wildlife authorities [M1; M2; W1].

The fire burned between 3360 – 4400 m above sea level, burning climax 'heather zone' forest, "spongy" [R1] Afroalpine moorland, and methane-rich bogs [M1], all with no recorded history of wildfires. (Fig. 3a; UNEP, 2022). Photographs from March 2012 show indicators of high burn severity (Fig. 3b), while images from July 2024, twelve years later, reveal regrowth limited to a maximum of 2.5 m, with the upper canopy still vacant (Fig. 3c). These slow growth rates and an observed scarcity of heather in the regrowth succession indicate that natural recovery will require several decades.

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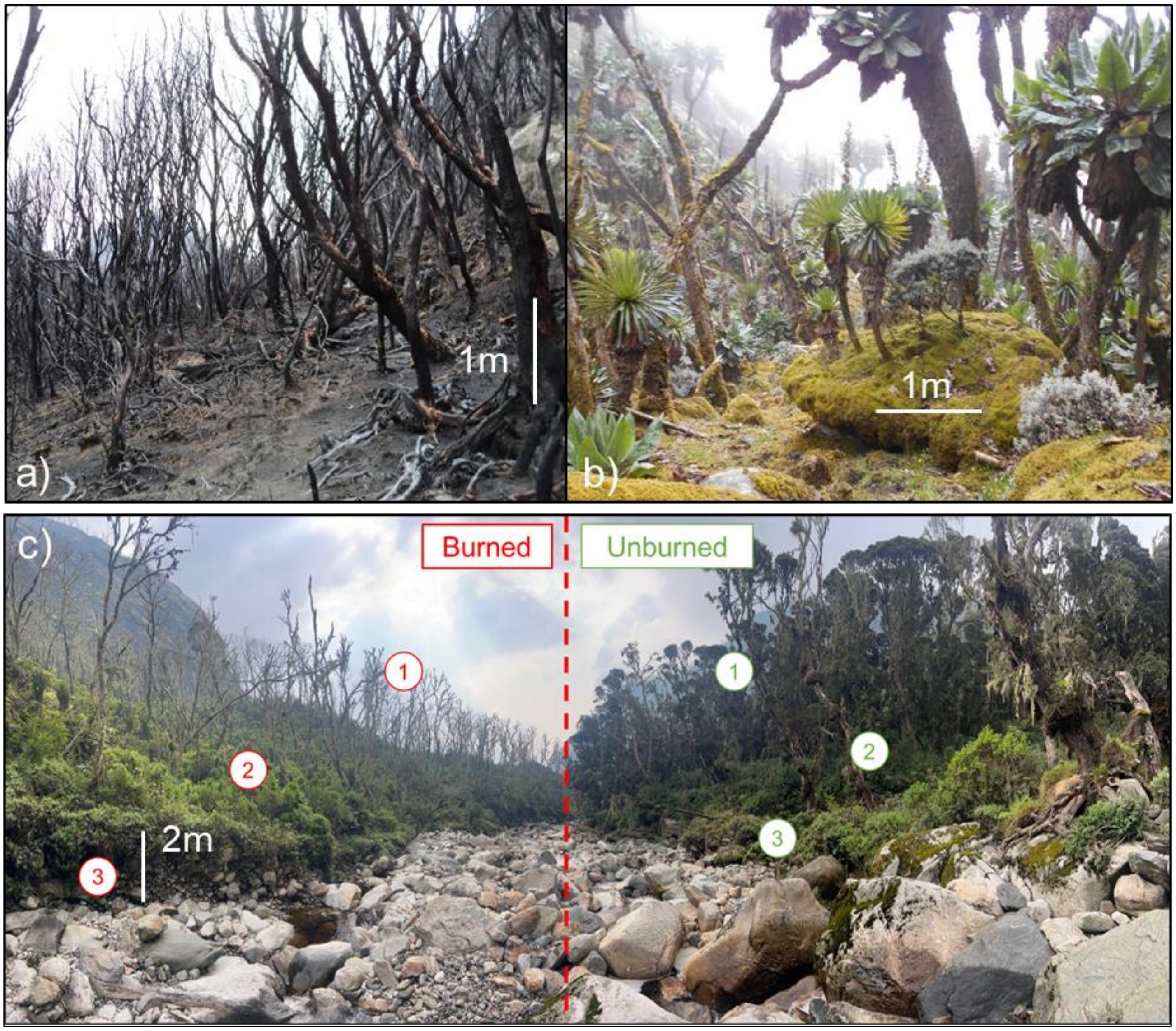

#	Attribute in c)	Burned	Unburned
1	Upper canopy	Dead ericaceous trees, vacant canopy	Mature ericaceous canopy
2	Lower canopy	Regrowth up to 2.5 m, heather scarce	Mature, dense vegetation
3	Riverbanks	Steep, unvegetated banks of exposed and unconsolidated glacial till	Sloped banks and coarse material anchored by vegetation

170 **Figure 3: a) burned ericaceous 'heather zone' vegetation 1-month after the wildfire in March 2012; b) Mature Afroalpine moorland vegetation prior to the wildfire (March 2011); c) upper course of the River Nyamwamba at 3380 m elevation in July 2024, where the river had acted as a firebreak to provide direct comparison between unburned and recovering burned sections of the ericaceous forest. Scale bars correspond to the tree trunk in 3a, the boulder in 3b and the river cliff in 3c. The associated table describes the ecological properties of the burned and unburned areas.**

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Wildfire-driven Interactions

The wildfire has a probability-increasing interaction with four secondary hazards. First, burning of soils and vegetation cover increased surface erosion and runoff to river channels, raising turbidity, carrying ash and peat, and introducing biological contaminants. Respondents recalled a strong smell "like methane" after the fire [M1; M2; R1; G1], highlighting wildfire-driven runoff pollution (#1). Second, reduced interception and infiltration capacity increased peak discharges at shorter lag times, driving a marked rise in fluvial flooding (#2; Sect. 3.2). Erosion has also accelerated due to higher discharges and loss of root cohesion (#3; Sect. 3.4), which, together with higher peak flows after the wildfire, enhanced the conditions for debris flow formation due to greater sediment supply [M1; R1] (#4).

Additional relationships where the wildfire has catalysed other hazard interactions are numerous, but evidence for these cannot fully be established without intensive monitoring and field experimentation. Based on hydrological theory, some interactions catalysed by the wildfire's effects would include:

- *River channel erosion-triggering-landslide (#8):* increased discharge after the wildfire catalyses the generation of landslides caused by erosive undercutting from higher river erosion rates
- *Landslide-increasing probability-river erosion (#12):* increased discharge catalyses the contribution of landslides to later erosion by transporting landslide talus and using the sediment as erosive tools for abrasion
- *Debris flood-triggering-river erosion (#14):* increased discharge catalyses erosion during debris flood events by increasing the erosive power of the flow
- *Landslide-increasing probability-fluvial flood (#16):* increased discharge increases the volume of water accumulating in damming and bursting flood mechanisms after landslides

Although many of the other hazards in the cascade are responsible for additional catalysing relations, we only present examples for the wildfire hazard in this study. This is to emphasise that the fire has not only increased the probability of four secondary natural hazards at the start of the cascade, but it is also catalysing subsequent interactions between other hazards.

3.2 Flooding

All twelve respondents reported heightened flood risk in the Nyamwamba catchment. Five attributed this directly to changes in hydrological processes caused by the 2012 wildfire [M1; M2; G1; G2; R1], while others cited land use change [N1; W1; R3], climate change [N1; I2], or the discontinuation of dredging [I1; R2]. A government official explained that "*the burning is the reason we are now having the floods annually… we know how useful wetland vegetation is in controlling floods, releasing water slowly*" [G1]. Similarly, a local guide described the flood-buffering role of the alpine wetlands: "*the moss was like a big*

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1 m thick sponge, it soaked up all the rain… 20 or 30 km² of rock that was once boulders covered in moss is now bare” [R1]. Table 2 documents ten flood events since 2012, all exceeding in intensity the two documented events during the preceding 12 years, with the 2013 and 2020 debris floods requiring international humanitarian appeals (Act Alliance, 2020; Delforge et al., 2025; Okiror, 2020).

Table 2: Timeline of flood events of the Nyamwamba River documented by humanitarian databases and grey literature since 2000. The dates of the two most intense flood events are highlighted bold.

Date	Area(s) Affected	Description & Impacts
1st May 2001	Rukoki, Kilembe	1 death and 300 people affected by flooding in Kasese District (Delforge et al., 2025; DesInventar, 2025)
8th September 2010	Rukoki, Ihandiro	A house, truck, pipeline and fields of crops destroyed by minor riverine flooding (Delforge et al., 2025).
February 2012 – Wildfire burns 30.75 km² of the Rwenzori National Park		
1st & 5th May 2013	Kilembe, Kasese District	Flooding in the Nyamwamba, Mubuku, Bulemba and Kitakena rivers displaced 25,445, with 13 deaths and US$4,055,000 of damage (Delforge et al., 2025). Formal humanitarian response appeal of $220,497 made by ACT Alliance (Act Alliance, 2013).
14th May 2014	Kasese town	3,725 affected and 4 deaths in Kasese (DesInventar, 2025).
18th June 2014	Kilembe	Flooded hospital and secondary school (Asiimwe, 2014).
18th April 2016	Kanamba, Kanaka, Kasese District	10,000 affected and an estimated $3,428,000 of damage following flooding of the Nyamwamba, Sebwe and Mubuku rivers between 4th – 18th April (Delforge et al., 2025; DesInventar, 2025; Juma, 2016).
4th July 2017	Kilembe	4 killed in the Kilembe Valley (DesInventar, 2025).
5th May 2020	Kasese District	173,000 people affected in 24,760 houses across Kasese and Bundibugyo Districts following flooding of major rivers (Delforge et al., 2025). River Nyamwamba overflow submerged the Kilembe Mines hospital, with over 1,200 people displaced in Kasese town (Act Alliance, 2020; Flood List News, 2020a, 2020b). Formal humanitarian appeal for assistance made by the Ugandan Red Cross to support the displaced (Okiror, 2020)
23rd May 2021	Kilembe	3 deaths and 134 affected following flooding and landslides in Kilembe town (Delforge et al., 2025).
18th May 2023	Kasese District	1,016 people affected, and 23 deaths recorded between 24th April and 18th May due to multiple floods of the Muhokya, Mubuku, Sebwe and Nyamwamba rivers (Delforge et al., 2025).
22nd May 2024	Kilembe, Kasese town	Sudden change of river course during high flow. Debris flows, riverine flooding and mudslides in the Nyamwamba catchment displaced 5,389 people in Kasese town (New Vision, 2024).
7th September 2024	Kasese Town	2 deaths and extensive damage to key infrastructure including schools, roads, bridges and 120 houses. Change of course of river during high flows breached same location as the 22nd May 2024 flood (ReliefWeb, 2024).

The wildfire has increased the frequency and magnitude of fluvial flooding, but also introduced two new mechanisms of flooding, with gravity-driven debris floods and avulsion floods linked to increased mass movement (Sect. 3.3) and erosion (Sect. 3.4) in the catchment [M1; M2].

Fluvial flooding

Vegetation and soil loss following the wildfire reduced interception, infiltration, and water retention capacity, amplifying the river's discharge response to rainfall. The fluvial flooding of unprecedented intensity on 5th May 2013 followed rainfall of only a 6.6-year estimated return period (Jacobs et al., 2016). Two respondents emphasise that a lack of lived experience prior to this first flood created additional vulnerability among affected communities: *"2013 - that was when we were all surprised. I could not believe what I saw"* [I1]; *"we were not prepared because we had never experienced such magnitude"* [M1]. Seven years later, an industrial worker recalled the 2020 event as "*an 800 cumecs flood… higher than our professional hydrologist's modelling of a 1000-year flood event*" [I2].

Debris flooding

Two floods (2013 and 2020) included debris flows, confirmed in video footage and respondent testimony [M1; R1]. A water authority described "*entire mahogany trees coming down as flood load*" [M1], while a resident noted "*moving rocks two times the size of a minibus*" [R1]. Field photos (Appendix G) confirm extensive boulder deposition on the floodplain, and the river has since shifted from a pre-wildfire meandering form with vegetated banks to a braided morphology laden with coarse crystalline sediment (Appendix K).

Avulsion flooding

Elevated erosion rates and sediment deposition have heightened the risk of avulsions [M1; M2]. On 22nd May 2024, for example, the Nyamwamba breached its outer bank upstream of Kasese town, inundating Kiwa hot springs and displacing 5,389 people [M1] (Table 2).

Flood-driven Interactions

High flows during fluvial and debris floods damage engineered flood defences, increasing their own probabilities of future breaches in self-perpetuating positive feedback (#5; #14; Appendix H). At the same time, their elevated velocities and turbulence generate shear stress and hydraulic action that trigger river erosion (#6; #15). GIS analysis confirms that the years of greatest erosion (2013 and 2020) coincided with the largest flood events [M1; R1] (Sect. 3.4).

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3.3 Landslides

Landslides caused by lateral river erosion undercutting riverbanks and hillslopes (Jacobs et al., 2016) have accelerated since the wildfire due to higher post-fire discharges and sediment loads [I1; M1; M2; R1]. In addition, the initial destruction and exposure of formerly stable riverbanks during the wildfire and 2013 flood has worked to further increase the probability of mass movement into the river [M1; G2]. Previously, graded banks of unconsolidated quaternary sediment were anchored by climax vegetation. Now, vertical river cliffs are exposed to direct erosion and undercutting at sites throughout the river's long profile. As one local government representative describes, "*when the floodwaters come down, they remove soil and grasses to expose more boulders, and then you will have a landslide*" [G2]. This process is visible in Fig. 3c, where the riverbanks on sections of the burned side are now steep, unvegetated cliffs of exposed and unconsolidated glacial till.

Landslide-driven Interactions

Landslides increase the probability of fluvial flooding by filling the channel with sediment and reducing the river's discharge capacity (#16). Five respondents have also witnessed a mechanism of temporary landslide damming and bursting "*in the space of a few minutes*" [M1] during high flow events, from which surges of sediment and discharge activate fluvial floods and debris flows [G1; M1; M2; R1; W1]. As one resident recalls: *"suddenly, I heard a roar like a plane taking off at Entebbe Airport. Two landslides cut off the river and created a dam behind it, then soon after there were entire trees pole vaulting over the debris"* [R1].

Five respondents describe landslides as being in a positive feedback process with erosion (through reciprocal interactions #17 and #8), whereby landslides add load to the river, accelerating lateral erosion by diverting flow to the riverbanks and causing further landslides [R1; M1; M2; I1; G2]. Landslides also trigger heavy metal pollution through the rotational slumping of solid Co-Cu tailings at Kilembe Mines into the River Nyamwamba (#18; Sect. 3.5).

3.4 Erosion

The cumulative annual eroded river channel area (Fig. 4a) shows a sustained increase in the river's rate of erosion by a factor of 2.64 following the 2012 wildfire, and the average middle-lower course channel width has increased sevenfold between 2010 – 2021, from 16.9 m to 123 m. Rapid erosion has destroyed agricultural land [M1; M2; G1; G2], residential property, and critical road infrastructure [M1].

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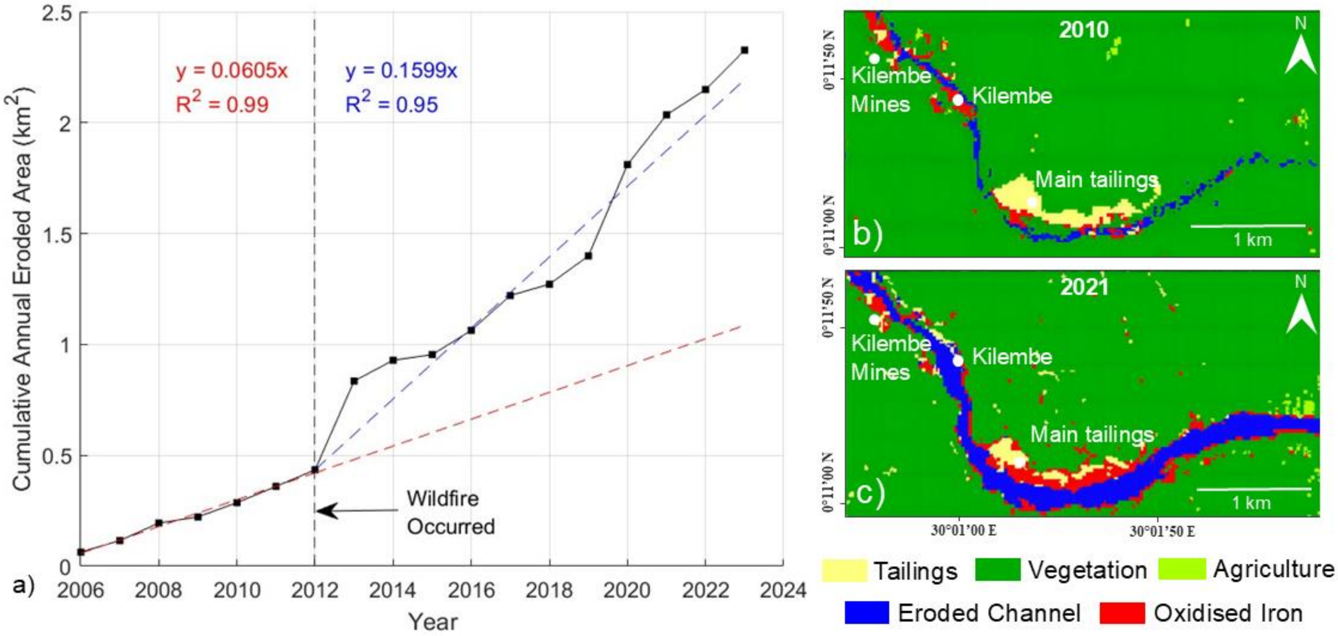

Figure 4: a) Annual cumulative eroded area in a 20 km mid-lower course section of the River Nyamwamba, calculated as the increase in eroded channel area between each year's supervised classification; b) supervised classification of a February 2010 Landsat-7 image; c) supervised classification of a February 2021 Landsat-8 image.

Erosion-driven Interactions

Since the wildfire, accelerated erosion has both shifted the river channel closer to populated areas of Kasese town and reduced its discharge capacity, thereby increasing the probability of urban flooding [R1; G1; G2; M2; I2; R3] (#7). Erosion also increases the probability of avulsion flooding, as exemplified by the May and September 2024 floods, by filling the channel with sediment bars that divert flow towards riverbanks [M1] (#12), whilst directly triggering avulsion floods when it breaks through unconsolidated banks [M1; M2] (#11). Contributions of sediment to the main Nyamwamba channel also increase the probability of debris floods [M1; R1] (#13).

Erosional undercutting destabilises slopes and directly triggers landslides (#8), consistent with Jacobs et al.'s (2016) mapping of 14 bank-failure slides during the May 2013 multi-hazard event. This lateral undercutting and exposure of vertical river cliffs is also described by three respondents as putting erosion in self-perpetuating positive feedback, by increasing the probability of further erosion at exposed banks [G1, G2, M1] (#9).

Channel widening breached the Kilembe mine copper-cobalt tailings deposit in 2014, since triggering heavy-metal pollution downstream that now presents a major risk to public health (#10; Sect. 3.5).

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3.5 Pollution

Immediately after the 2012 wildfire, community members reported increased turbidity and a smell "*like methane*" [M1] in the river. This is still reported during high discharge twelve years later, which four respondents believe to be due to runoff (non-point source) pollution through exposed bogs and organic-rich glacial sediments in the fire-affected and eroding upper catchment [M1; M2; R1; G1] (#1).

Beyond this diffuse pollution, accelerated river erosion (#10) and landslides (#18) have inputted an estimated 744,000 tonnes of a 15 Mt Kilembe Mines Co-Cu tailings deposit directly into the River Nyamwamba (mapped in Appendix K). Satellite imagery and field photographs show erosional cliffs, slump scars and new channels within tailings areas, and evidence of acid mine drainage from distinctive iron oxide precipitation (Fig. 3c; Appendix E). Elevated Co, Ni, Cu, Fe, Al, S, Zn, As, Cd and Mn river contamination has previously been attributed to leaching of the Co-Cu mine (Abraham & Susan, 2017; Mwesigye et al., 2016; Mwesigye & Lawrence, 2024; Mwongyera et al., 2014).

Five respondents identified this solid waste pollution as a major concern for public health [M2; W1; G1; G2; R1]. The river is used by 38% of its adjacent population for drinking, and by many more indirectly through crop-irrigation and groundwater abstraction (Abraham and Susan, 2017; Mukisa et al., 2020). In addition to waterborne risks, long-term contamination of arable soils by deposited mine waste raises concern for food safety [M1; M2; G1; G2]. As one Ministry of Water official noted, *"in Kasese District, their teeth are turning brown with yellow patches, and we have been told that many people in this region are ailing with cancer"* [M2]. Local environmental managers also expressed concern for downstream ecosystems in Queen Elizabeth National Park and Lake George, where protected flora and fauna may be affected by the pollution and vegetation dieback observed in Kasese town [G2; M1, R1, W1].

4 Discussion: Implications for Management

The intensity and persistence of the Rwenzori hazard cascade highlights how wildfires in mature, fire-sensitive mountain ecosystems can impose long-lasting risks on downstream communities. Unlike fire-adapted systems where vegetation recovers quickly to dampen hazard impacts, recovery in these environments is slow, and positive feedback mechanisms sustain elevated risk. By characterising hazard interactions in full, this study identifies entry points for intervention. Management approaches that systematically impede hazard interactions can help unravel cascades (Gill and Malamud, 2016).

The principal way to impede this cascade is at the top (interactions #1-4), by promoting ecosystem recovery and attenuating the elevated runoff and river discharge driving other hazards. In the Rwenzori, authorities implemented a mix of hard engineering, community-centred and nature-based solutions that has saved lives (see Appendix L). However, the prevailing approach to wildfire restoration has been to await natural recovery. This passivity missed a critical window to implement soil stabilisation and runoff attenuation solutions such as mulching, contour felling and forest restoration (Papaioannou et al., 2023; Robichaud et al., 2013; Scheper et al., 2021), and allowed lower canopy vegetation to establish ahead of upper canopy tree species in the ericaceous zone (Fig. 3c). The challenge now is to develop recovery and discharge attenuation solutions in a partially recovered ecosystem. Addressing this requires post-wildfire expert assessment to guide restoration planning and build an evidence-base for financing solutions (Veness and Buytaert, 2025).

In the later stages of the Rwenzori cascade, erosion emerges as a key driver of multiple hazard interactions and positive feedback processes. It has accelerated landslides, amplified debris flows, triggered flooding, and caused a major water pollution hazard now requiring urgent investigation of its scale and health impacts. Stabilising riverbanks is a critical intervention to mitigate erosion and therefore impede its cascading interactions. We recommend integrating existing dredging, levee construction, and nature-based approaches to achieve this (Appendix L; MoWE, 2022). In particular, repositioning coarse sediment to riverbanks can help protect eroding river cliffs, regrade unstable slopes, and create conditions for in-channel vegetation to anchor finer sediments and restore soil, thus mimicking the stable, unburned riverbank morphology seen in Fig. 3c (Sanches Fernandes et al., 2020). These measures are urgent in the mid-catchment to protect communities and limit further mobilisation of solid mine waste, but also advisable in the upper catchment to reduce sediment generation and landslide risk. Developing an effective approach to bank stabilisation in the Rwenzori could serve as a blueprint for similar future events.

Montane environments globally, especially those without a history of fire, require greater investment in monitoring and research into post-wildfire hazard cascades (Arango Carmona et al. 2025; Wimberly et al. 2024). The lack of comparable case studies makes it difficult to determine whether the Rwenzori represents an outlier or part of a broader emerging trend. However, the intensity of the Rwenzori cascade, following a burn area of just 31 km^2, is a warning to trigger post-fire risk assessments at lower thresholds of burn area and severity when the fire occurs in a fire sensitive mountain ecosystem. Expanding research in similar regions will help build an evidence base of common cascading interactions and best practices for their management.

5 Conclusions

This study has characterised a post-wildfire multi-hazard cascade in a tropical montane catchment, demonstrating how the burning of a pristine, fire-sensitive mountain ecosystem can initiate cascading hazards of exceptional intensity and persistence. As fire regimes continue to shift to higher altitudes under climate change, there is an emerging risk of similar hazard cascades for downstream communities in tropical mountain catchments worldwide.

In Uganda's Rwenzori National Park, in the twelve years after a 2012 wildfire burned 31 km^2 of mature forest and peatland, ten major floods with fluvial, debris or avulsion mechanisms occurred, with two debris floods requiring large-scale humanitarian responses. Increased river discharge after the fire caused a 2.64-fold increase in erosion rates and increased the

https://doi.org/10.5194/egusphere-2025-5106
Preprint. Discussion started: 24 November 2025

probability of landslides, which have together driven a sevenfold increase in river channel width over nine years. Urban and agricultural areas now face a real-time risk to public health due to the erosion and mass movement of 744,000 tonnes of copper-cobalt solid tailings into the River Nyamwamba since 2014. This discrete escalation of hazards, interactions and impacts is sustained by the slow recovery of vegetation poorly adapted to fire regimes, and multiple positive feedbacks between hazard interactions.

The Rwenzori case highlights a need to recognise post-wildfire hazard cascades as a long-term risk in tropical mountain environments, especially in newly fire-prone areas with no prior history of burning. We recommend post-fire risk assessments and research, even for relatively small burn areas, when future fires occur in previously unburned or fire sensitive mountain ecosystems. A better understanding of interactions between hazards identifies intervention points, where interactions can be impeded through early actions that prevent ecosystem impacts from becoming entrenched long-term. To this end, remediation of the burned zone should always be a priority to accelerate ecosystem recovery and attenuate elevated runoff.

More monitoring and research of global case studies is needed to establish the prevalence and intensity of tropical mountain wildfire hazard cascades, and best practices for their management. This study has additionally underscored the value of integrating qualitative data and local knowledge into such studies. Interviews were critical to identifying key hazard interactions that would not have been captured through physical or remote data alone. Interdisciplinary research, through close partnerships between academic and local stakeholders, can improve collective visibility on this emerging climate risk and accelerate the development of shared solutions.

6 Data Availability

The interview data is confidential according to ethical and data sharing restrictions. The GIS files are available on GitHub (https://github.com/will-veness/wildfires-uganda) and will be available in Zenodo upon full publication.

7 Competing Interests

We declare no competing interests.

8 Author Contributions

William Veness: Writing – review & editing, Supervision, Writing – original draft, Visualization, Validation, Software, Resources, Project administration, Methodology, Investigation, Funding acquisition, Formal analysis, Data curation, Conceptualization

Martha Day: Writing – review & editing, Writing – original draft, Visualization, Validation, Software, Resources, Methodology, Investigation, Formal analysis, Data curation, Conceptualization

Anthony C. Ross: Writing – review & editing, Investigation, Data Curation, Validation

Yazidhi Bamutaze: Writing – review & editing, Funding acquisition, Data curation, Supervision, Investigation, Validation, Conceptualization.

Jiayuan Han: Methodology, Investigation, Data Curation, Visualiazation, Software

Douglas Mulangwa: Project administration, Resources, Data curation, Investigation

Anthony Mwesigwa: Writing – review & editing, Project administration, Data curation, Investigation

Emmanuel Ntale: Project administration, Resources, Investigation

Callist Tindimugaya: Project administration, Resources, Methodology, Supervision, Conceptualization

Brian Guma: Project administration, Resources, Methodology, Investigation, Conceptualization, Supervision, Validation

Elisabeth Stephens: Writing – review & editing, Investigation, Project administration, Resources, Data Curation, Investigation, Validation

Wouter Buytaert: Writing – review & editing, Supervision, Methodology, Investigation, Project administration, Conceptualization, Validation

9 Acknowledgements

We thank the Uganda Red Cross Society, the Ministry of Water and Environment, Kasese Municipality, and all local collaborators and community members in Kasese District for their contributions to this research. We are grateful to the Uganda Wildlife Authority and Rwenzori Trekking Services for facilitating access to the Rwenzori Mountains National Park. This study was initiated through support from Imperial College London's African Research Universities Alliance (ARUA) partnership, which enabled collaboration with Makerere University. The Red Cross Red Crescent Climate Centre is also gratefully acknowledged for its role in establishing the partnership with the Uganda Red Cross and the Ministry of Water and Environment.

https://doi.org/10.5194/egusphere-2025-5106
Preprint. Discussion started: 24 November 2025

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https://doi.org/10.5194/egusphere-2025-5106
Preprint. Discussion started: 24 November 2025

11 Appendices

Appendix A: 1-Month Standardised Precipitation Index Calculation for January 2012

ERA5 monthly averaged reanalysis total precipitation data was downloaded from 1974 – 2024 for the pixel covering to the burned area (centroid coordinates: 0.4°N, 29.8°E; Copernicus Climate Change Service (C3S), 2017). This was processed in MATLAB following McKee et al.'s (1993) method to determine the monthly-SPI for January 2012.

Appendix B: Eroded Tailings Volume Calculation

The average original height of the tailings was calculated to be 23 m, assumed to be flat across the original dammed area, which was calculated (32945 m^2) from historic satellite imagery.

This average height (23 m) was multiplied by the eroded footprint area (m^2) to get a volume, then volumetric subtractions were made to account for the originally sloped (55 degrees) walls of the tailings dam and the wedges of slumped material yet to be eroded at the foot of the collapsed tailings escarpments. The volume of these wedges was calculated from the slope angle and height of their triangular cross-section, multiplied by their width parallel to the eroded tailings escarpment.

The tonnage of eroded tailings was then calculated by multiplying their estimated volume by their assumed average dry density (1.5 t/m^3) based on standard values for copper-cobalt tailings (Williams, 2015).

Appendix C: Semi-Structured Interview Template

Background

- What organisation do you represent?
- What is your role?
- What is your experience of hazards in the Rwenzori?

Perceptions of changing hazard risk

- Do you feel the risk of hazards have changed (in the Nyamwamba catchment)? How?
- If yes, why do you feel risk is changing?
- Do you feel the river Nyamwamba/Mubuku/other rivers have changed?
- If yes, why do you think this change has happened?

Awareness and efficacy of existing management strategies

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- What existing strategies are in place to manage hazard risk in the Rwenzori?
- Do you feel these strategies are working?

Potential alternative management strategies

- What strategies do you feel would better reduce hazard risk in the Rwenzori?
- Why do you think these have not been implemented yet?
- Do you feel nature-based solutions could be used to manage these hazards?

Appendix D: Evidence for Multi-Hazard Cascade Interactions

Table D1: Description of the hazard interactions in Fig. 2 and supporting evidence.

#	Initiating Hazard	Affected Hazard	Interaction Description	Evidence
1	Wildfire	Runoff pollution	Increased probability. Burning of soils and vegetation cover increased their erosion and runoff to the river channel. This hazard is also catalysed by higher rates of erosion increasing the delivery of soil, ash and peat to the river.	Four interview respondents describing increased turbidity immediately after the wildfire and during high flows, with a smell “like methane” [M1, M2, R1, G1].
2	Wildfire	Fluvial flood	Increased probability. Burning of vegetation has reduced interception and root uptake of precipitation, increasing surface runoff to the channel. This has increased peak discharges at reduced lag times following peak rainfall events. The burning and erosion of mature soils has also reduced their infiltration and storage capacities, therefore increasing runoff.	Humanitarian data of 10 flood events since 2012 exceeding the impacts of any documented flood in the 12 years prior (Table 1). Interviewee accounts [M1, M2, G1, G2, R1], e.g. “*the burning is the reason we are now having the floods annually… we know how useful wetland vegetation is in controlling floods, releasing water slowly*” [G1].
3	Wildfire	River erosion	Increased probability. Wildfire’s burning of vegetation and erosion of soil has increased runoff, peak river discharge, and therefore the erosive power of the river. Initial erosion and mass movement also exposed river cliffs, which is increasing the probability of (and catalysing) further erosion in a positive feedback process.	GIS analysis calculating an erosion rate increase by a factor of 2.64 due to the wildfire (Fig. 4). Photographs of exposed river cliffs within wildfire affected areas (Fig. 3c and Appendix F).
4	Wildfire	Debris flood	Increased probability. Wildfire has increased peak river discharge by the burning of vegetation and soil which modulate discharge. It has also increased sediment generation through augmented erosion and mass movement, improving the conditions for debris flow development.	Two interview respondents explain and show camera footage of 2013 and 2020 debris flows, described as unprecedented before the fire [M1, R1] (Table 1). Jacob’s et al.’s (2016) reconstruction of debris flows during the May 2013 flood. Field photographs of boulder deposition on the delta and distal flood plain (Appendix G).

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5	Fluvial flood	Fluvial flood	Increased probability. Fluvial floods damage engineered flood defences, increasing the probability of future breaches.	Photos of damaged flood defences (Appendix H).
6	Fluvial flood	River erosion	Triggering. Higher flow velocities and turbulence during fluvial floods exert shear stress, abrasion and hydraulic action to erode river banks.	GIS analysis shows the years of highest erosion occurred in 2013 and 2020, the years of the largest debris and fluvial floods [M1, R1] (Fig. 4**)**.
7	River erosion	Fluvial flood	Increased probability. Eroded material fills and reduces the channel's carrying capacity for discharge. Erosion has also relocated active channels closer to residential areas.	Change in river morphology to a sediment-laden braided system indicating increased deposition and channel switching (Fig. 4b).
8	River erosion	Landslide	Triggering. Lateral and vertical erosion of riverbanks undercuts and destabilises hillslopes, increasing local shear stresses to failure.	Jacobs et al. (2016) map 14 landslides triggered by scour and bank failure from river erosion.
9	River erosion	River erosion	Increased probability. Erosion of banks exposes steep, unstable river cliffs to further erosion.	Photos of erosional river cliffs incising into hillslopes at multiple sites (Appendix F). Interviewee descriptions [G1, G2, M1]
10	River erosion	Heavy metal pollution	Triggering. River erosion has breached the main 15 Mt solid Co-Cu Kilembe Mines tailings deposit and other smaller deposits within the town.	Satellite images and field photographs (Fig. 4) show erosive river cliffs and new channels within the original tailings area. Field observations of downstream deposition of tailings and iron precipitates (Appendix E). Four respondents consider waste deposition a major concern for public health and a potential cause of vegetation death on the riverbanks [M2, W1, G2, R1].
11	River erosion	Avulsion flood	Triggering. Increased erosion of river banks causes channel-switching and subsequent avulsion floods.	Humanitarian data and interview respondents [M1, M2] describing the 22nd May 2024 avulsion flood impacting Kasese town (Table 1).
12	River erosion	Avulsion flood	Increased probability. Higher rates of upstream erosion increase downstream deposition in channel bars, diverting flow towards river banks.	Interview respondents [M1, M2] describing the 22nd May 2024 avulsion flood impacting Kasese town (Table 1) and the increased rate of deposition that has raised dredging and channel clearance costs since the 2012 wildfire [M1, M2, R1].
13	River erosion	Debris flood	Increased probability. Erosion provides additional sediment that improves the probability of debris flow formation.	GIS analysis of increased channel area and width (Fig. 3) filled with coarse sediment in a braided system (Appendix K). Two respondents describe debris flows as unprecedented before the fire [M1, R1].
14	Debris flood	Debris flood	Increased probability. Debris floods damage engineered flood defences, increasing the probability of future breaches.	Photos of damaged flood defences (Appendix H).

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15	Debris flood	River erosion	Triggering. Debris flows have high erosive power (Church & Jakob, 2020).	GIS analysis shows the years of highest erosion occurred in 2013 and 2020, the years of the largest debris and fluvial floods [M1, R1] (Appendix K**)**.
16	Landslide	Fluvial flood	Increased probability. Landslide material fills and reduces the channel's carrying capacity for discharge. Landslides also increase the probability of fluvial (and debris) flooding through temporary damming and bursting mechanisms that create surges of discharge.	Jacobs et al. (2016) mapped 29 landslides during the May 2013 flood that directly entered the River Nyamwamba. Five respondents describe a mechanism of temporary landslide damming and bursting "in the space of a few minutes" [M1] during peak rainfall events in the upper-catchment [G1, M1, M2, R1, W1].
17	Landslide	River erosion	Increased probability. Landslides increase sediment load and the subsequent erosive power of the river through abrasion.	Field photographs of slump scars on river banks (Appendix F, Appendix H). Analysis by Jacobs et al. (2016) showing landslides directly entering the river system.
18	Landslide	Heavy metal pollution	Triggering. Rotational slumping of the soft, unconsolidated tailings into the River Nyamwamba causes heavy metal contamination of water and sediment.	Satellite images and field observations (Appendix K) show rotational slump scars throughout the affected tailings. Four respondents consider waste deposition a major concern for public health and a potential cause of vegetation death on the riverbanks [M2, W1, G2, R1].

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Appendix E: Tailings Pollution Photographs

Figure E1: Acid mine drainage at location 0.18599N, 30.01951E, 25 July 2024.

https://doi.org/10.5194/egusphere-2025-5106
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710 **Figure E2: Acid mine drainage at location 0.19879N, 30.01139E, 3 August 2024.**

https://doi.org/10.5194/egusphere-2025-5106
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Figure E3: Tailings sedimentation in the Nyamwamba channel, 0.18652N, 30.01986E, 25 July 2024.

https://doi.org/10.5194/egusphere-2025-5106
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Appendix F: Exposed River Cliffs Photographs

715 **Figure F1: River cliff exposure at 0.29291N, 29.93596E – 28 July 2024.**

https://doi.org/10.5194/egusphere-2025-5106
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Figure F2: River cliff style erosion of house foundations in Kilembe, 0.20603N, 30.00822E – 24 July 2024.

https://doi.org/10.5194/egusphere-2025-5106
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Figure F3: River cliff at 0.23742N, 23.97568E – 1 August 2024.

https://doi.org/10.5194/egusphere-2025-5106
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Figure F4: River cliff at 0.23715N, 29.97601E – 1 August 2024.

https://doi.org/10.5194/egusphere-2025-5106
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Appendix G: Flood Plain Boulder Deposition Photographs

Figure G1: 0.20285N, 30.00908E - 7 June 2023.

https://doi.org/10.5194/egusphere-2025-5106
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Figure G2: 0.19528N, 30.01544E - 25 July 2024.

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https://doi.org/10.5194/egusphere-2025-5106
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Appendix H: Damaged Flood Defences Photographs

Figure H1: 0.18981N, 30.07408E, 26 July 2024 (damaged bamboo nature-based solution).

https://doi.org/10.5194/egusphere-2025-5106
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Figure H2: 0.21387N, 30.00558E, 7 June 2023 (damaged gabions).

https://doi.org/10.5194/egusphere-2025-5106
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Appendix I: Landslide Talus Entering the River Photographs

Figure I1: 0.29291N, 29.93596E – 28 July 2024.

https://doi.org/10.5194/egusphere-2025-5106
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Figure I2: Landslide scar at 0.23758N, 29.97570E – 1 August 2024.

https://doi.org/10.5194/egusphere-2025-5106
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Appendix J: Lower Course Deposition of Solid Mine Tailings Photographs

Figure J1: Deposition and acid mine drainage downstream of Kilembe Mines 0.19385N, 30.082355 E, 26 July 2024.

https://doi.org/10.5194/egusphere-2025-5106
Preprint. Discussion started: 24 November 2025

765 **Figure J2: Acid mine drainage from deposited solid tailings at location adjacent to Kasese town, 26 July 2024.**

https://doi.org/10.5194/egusphere-2025-5106
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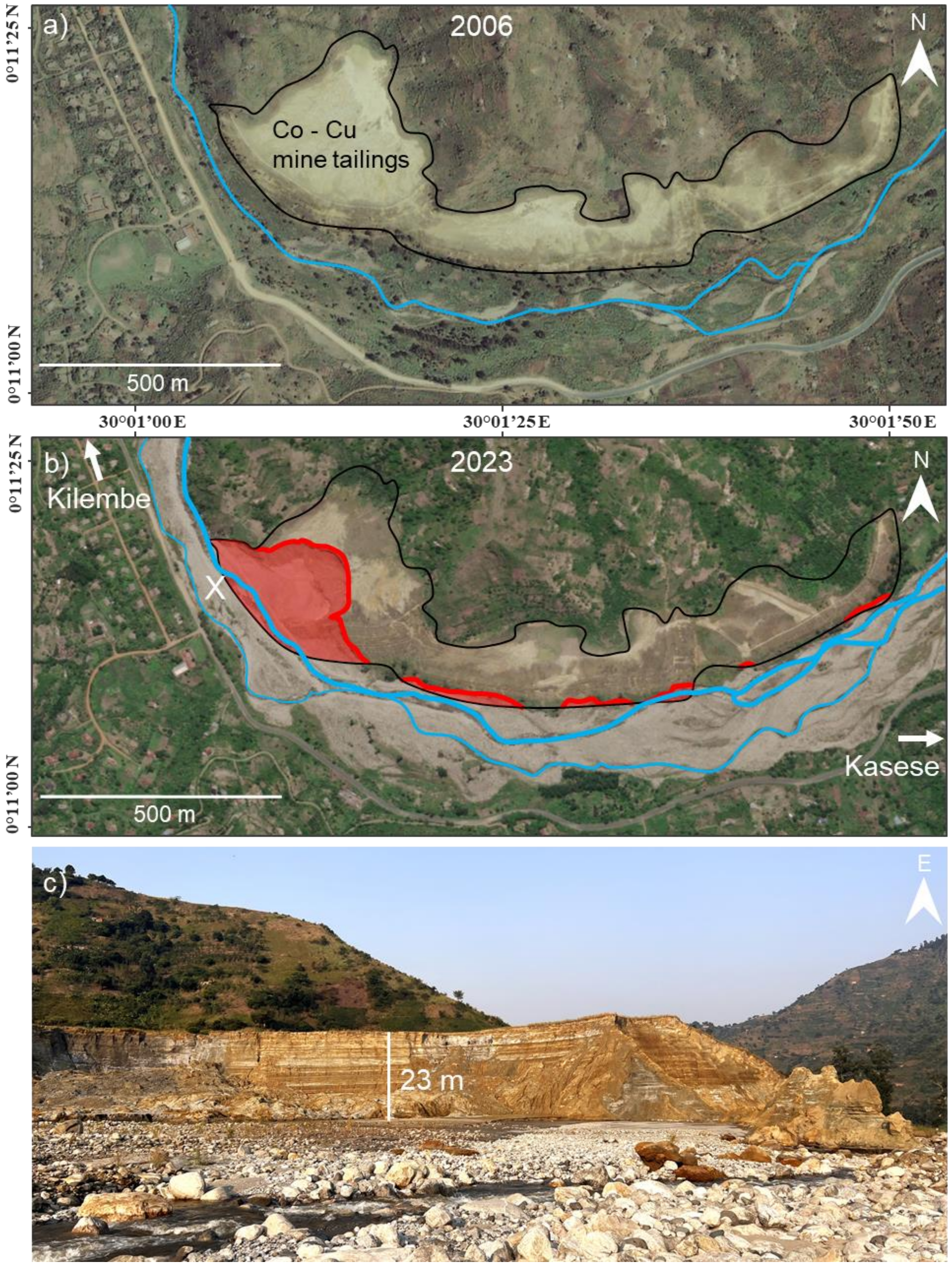
Appendix K: Kilembe Mines Co-Cu Tailings Erosion



770 **Figure K1 – a) Kilembe Mines tailings on 24th March 2006 (Maxar Technologies, 2024a); b) the same location on 10th Apr 2023 (Maxar Technologies, 2024b) where an estimated 744,000 tonnes of solid waste have eroded into the river. The black polygon outlines the original surface area of mine tailings, and the red polygon shows the area partly or fully eroded; c) photograph of a section of the eroded tailings taken in July 2024 at position X (b), facing east towards Kasese town.**

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Appendix L: Management Strategies and their Evaluation

The disaster risk management (DRM) strategies in Table L1 have been implemented in the Nyamwamba catchment since the May 2013 floods. Whilst relocation of communities experiencing near-annual flooding is considered desirable for mitigating their flood risk [M1, M2, N1], residents have opposed relocation due to existing community and land ties, lower living costs on flood plain and a lack of economic opportunity in areas proposed for relocation [M1, G2, N1]. Instead, therefore, strategies have focussed on protecting existing communities and informal settlements on the flood plain with hard engineering, community-centred and nature-based solutions (Table L1).

Table L1 – Summary of disaster risk reduction strategies in Kasese Districts observed during field reconnaissance and described by interview respondents.

Strategy		Description	Evaluation
Hard Engineering	Gabions (Figure L1a)	Installed in phases along a 2-3 km alongside Kilembe town to mitigate flooding and river channel switching.	"The Gabions have failed, they're very weak" [R1]. Damaged by minor flood events (Appendix H), and they have failed to prevent channel overtopping into Kilembe Town during the 22nd May 2024 flood [M2].
	Channelisation (Figure L1b)	A short 500 m channelised section of channel downstream from Kilembe, using concrete to ensure the stability of the road-bridge providing the only access route to Kilembe town.	"what they have done [at the bridge] is perfect…the narrow section never gets clogged up so the rocks pass through" [R1]. The solution has been positively received [R1, R2], but it is considered expensive, and river may switch channels if extended [M2].
	Dredging	5 km of channel is desilted (boulders are broken down and removed from the active water channels to the banks) in an irregular regime, typically funded after major floods such as those in 2013 and 2020 [M1, M2].	It costs US$4.5 million to clear 5 km of the channel and it needs to be performed annually to maintain a cleared channel [M2]. Residents recall successful desilting by a Canadian mining company until the 1971, so it is a positively viewed activity [G1, G2, W1] but may not be economically sustainable with the currently increased sediment flux of the river [M1]. It does not take place far enough upstream of Kilembe where debris flows generate [R1, M2].
Community-Centred Solutions	Flood Early Warning Systems (Figure L1c)	Communities of Kilembe, Kasese and Mubuku given early warnings through the Ugandan Red Cross and the Ugandan Ministry of Water and Environment (MoWE) following alerts of high rainfall.	"with early warning systems, less people are dying… people are more informed with better risk communication" [N1]. However, difficulties monitoring water levels due to high sediment loads and channel switching leaves early warning dependent on rainfall forecasts that are low-confidence in a convectional mountainous region [M2]. Expansion requires greater hydrological monitoring for more

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			accurate, confident and timely warnings [N1, M2].
	Resident Relocation	Relocation of displaced households from Kilembe and river banks to the Kasese lowlands, using emergency response funding following major 2013 and 2020 flood events [N1, M1]. Matched with investment to support alternative livelihoods independent of the river such as bee keeping [M1].	In most cases, residents have refused to relocate and are building informal homes [M1]. There is a need for expansion of livelihood incentives and longer-term support investments for their setup [G2]. High flood risk areas offer low-cost land, economic opportunities, free water from the river and many have attachment to lands from family history, mountain livelihoods and lived experience [N1].
	Participatory Desilting	Pilot project training individuals to convert river boulders into crafts, such as granite wash-basins to be sold to safari lodges and tourists.	It is not being completed at a scale that significantly impacts flood risk [M1], but it has been shown to successfully supplement family incomes (Ugandan Ministry of Water and Environment (MoWE), 2022). It requires longer-term investment and a plan for expansion and greater access to the market [N1, M1].
Nature-based Solutions (NbS)	River Bank Stabilisation (Figure L1d)	As part of a 2021 World Bank funded project (MoWE, 2022), a 10 km length of the Nyamwamba river banks have been planted with 30 m thick vegetation buffers to mitigate further lateral erosion. Seedlings planted included 35,000 Asper bamboo, 2,000 mango and 4,000 Mahogany, situated within a fenced zone to deter trespassing, logging, theft and interference by animals (MoWE, 2022).	An existing pilot in Mubuku has demonstrated 20 years of successful bank stabilisation [M1], however 2021 Nyamwamba planting has faced challenges of droughts, floods, termites [W1], death of seedlings due to heavy metal contamination by mine tailings, logging, and reluctant participation by some land owners. Rapid initial growth in patches require long term monitoring and evaluation, but bamboo planting is perceived as the most promising solution for landslide and erosion mitigation in the wildfire-affected zone and around the mine tailings [R1, M1, M2, I1, I2, W1]
	Soil and Water Conservation	Awareness raised among 1,420 land owners of methods available to reduce soil erosion and runoff. 750 were trained to implement the intervention and provided equipment, with 211 hectares of land modified by the addition of trenches and hedges in 2021-2022 (MoWE, 2022). Households encouraged to harvest rainwater instead of drinking from the river.	"there was actually a gentleman that implemented it on his own land, without us telling him to." Need for more land-owner co-operatives to share trainings, to share risk of failed implementation following land conversion, and to share tedious workloads [G2]. Rainwater harvesting reduces runoff, soil erosion on small plots and decreases heavy metal consumption from river water [M2, G2].
	Afforestation and Regrading of Hillslopes	825 hectares afforested through reforestation and agroforestry in the mid-catchment to reduce landslides, soil erosion and runoff to the river (MoWE, 2022).	Soil-water conservation trenches and soil-stabilising species increased coffee yields [F1]. Some respondents criticised soil-water conservation and afforestation efforts for focussing on the mid-catchment, when "99%"

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			of the sediment and discharge generation is taking place in the burned national park area upstream [R1, I2]. "Until we stabilise those areas [upstream] we will have these problems" [I2].

For hard engineering strategies, respondents believe that gabions are too weak to sustainably channelise the river [R1, M2] (Figure L1a), whereas there is demand for the successful concrete channelisation to be extended beyond Kilembe town centre [R1, R2, M2] (Figure L1b). Channel dredging is perceived to be a critical activity, not because of successful implementations since 2013, but due to successful historic programmes of dredging by mining companies when Kilembe mines was operational in the 1960s [G1, G2, W1, R1, M2]. For all hard engineering approaches, there is concern of an unsustainably high cost of maintenance, given the elevated rate of discharge, erosion and sediment generation in the Nyamwamba river [M1, M2].

Flood early warning systems piloted in Kilembe and Kasese using 2 local rain gauges and water level sensors have faced challenges of continuous automated data collection in hard-to-reach upstream locations, however, sharing of information between authorities and community representatives via Whatsapp has successfully coordinated evacuations following high flows and rapid dispatches of emergency respondents [N1]. A 2023 installation of a camera 5 km upstream of Kilembe, capable of international photo and video transmission at 1-minute intervals (Figure L1c), is considered a useful supplementary dataset for a more detailed interpretation by those with lived experience and indigenous knowledge of the river [N1, M2]. For rivers with a debris-flow model of flooding, setting qualitative thresholds of perceived flood severity from imagery may have more local predictive value than water levels in channels where channel location and roughness change frequently [M2].

A project funded by the World Bank and implemented by the Ugandan Ministry of Water and Environment (MoWE) in 2021 – 2022 has installed a range of nature-based (NbS) and community-centred solutions (MoWE, 2022). The NbS of river bank stabilisation in Kasese is considered especially promising [R1, R3, W1, M1, M2], using 35,000 asper bamboo seedlings and other economic crops in buffer zones on the mid-catchment river banks to prevent erosion. Despite challenges with drought, flooding, termites and metal-contaminated soils during the early implementation [W1, M2, G2], a previous project successfully stabilising the Mubuku river banks for 20 years [M1] and observations of stable bamboo forests in the upper catchment [R1] provide optimism for the project. Respondents are more critical of other parts of the project, including soil-water conservation and participatory desilting of the river (Table L1), for focussing on the mid-catchment around Kasese town, when discharge and sediment generation is taking place higher in the mountains [R1, I2].

“the assumptions made are well off beat; "99% of the water is coming from the park” – R1

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815 **Figure L1 – Photographs taken during June 2023 field reconnaissance: a) collapsed gabions adjacent to Kilembe town (for scale: 8 m channel width); b) channelisation using concrete embankments in Kilembe town centre (10 m channel width under bridge); c) photo from a camera transmitting photos at 1-minute intervals 5 km upstream of Kilembe town centre for flood early warning; d) river bank stabilisation adjacent to Kasese town including asper bamboo (4 m fencepost spacing).**

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Notably, there have been no DRR interventions so far in the wildfire affected area of the upper catchment, and no active mitigation of mine tailing erosion into the river Nyamwamba. In both cases, a low awareness of their impacts has inhibited action [M2, R1, W1, G2]. 7 of 12 interview respondents did not mention the 2012 wildfire when asked to describe factors affecting local flood risk, and only one small-scale academic study has assessed water quality in the Nyamwamba since large-scale erosion began in 2015 (Mukisa et al., 2020). Of the respondents aware of the wildfire [R1, W1, M1, M2, G2] and water quality problems [M1, M2, G1, G2] in the Nyamwamba catchment, all recommend restoration of the wildfire-affected area and urgent mitigation of further erosion into the river.