

1 **Temperature–RH Dependent Viscosity of Organic Aerosols**  
2 **from 273 to 303 K: Implications for Global N<sub>2</sub>O<sub>5</sub> Uptake**

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10 Summary: 21 pages, 11 figures, 1 table

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## 12 S1. Flow-cell calibration and particle conditioning time determination

13  
14 In both bead-mobility and poke-and-flow experiments, the flow-cell was calibrated for relative humidity (RH)  
15 using the deliquescence points of three salts including  $K_2CO_3$ , NaCl, and  $(NH_4)_2SO_4$  (Winston and Bates, 1960).  
16 The experimentally determined deliquescence RH at each measurement temperature 273, 283, 293, and 303 K were  
17 compared with literature. The RH accuracy was within  $\pm 2.0\%$  and the temperature was controlled within  $\pm 1$  K  
18 of the target.

19 Prior to each experiment, particles deposited on a hydrophobic substrate were conditioned at the target RH of  
20 the carrier gas to allow sufficient equilibration with the surrounding gases. Conditioning times were adjusted  
21 depending on the experimental temperature and technique. To evaluate whether the samples reached near-  
22 equilibrium with the target RH, we applied a method previously reported in the literature (Kiland et al., 2023a;  
23 Smith et al., 2021a; Kiland et al., 2023b; Maclean et al., 2021; Evoy et al., 2021; Smith et al., 2021b). This method  
24 involves comparing experimental conditioning time to the characteristic mixing time of water within organic  
25 aerosol (OA),  $\tau_{mix,H_2O}$ , calculated:

$$\tau_{mix,H_2O} = \frac{d_p^2}{4\pi^2 D_{H_2O}(T, RH)} \quad (S1.1)$$

26 Here,  $d_p$  represents the particle diameter.  $D_{H_2O}(T, RH)$  is the RH- and temperature-dependent diffusion coefficient  
27 of water in OA. The value of  $D_{H_2O}$  was calculated using the fractional Stokes–Einstein equation, which accounts  
28 for the link between viscosity and diffusion in cases where the diffusing species are comparable in size to, or smaller  
29 than, the molecules forming the matrix (Evoy, 2020):

$$D_{H_2O}(T, RH) = D_{H_2O}^\circ(T) \times \left( \frac{\eta_{H_2O}^\circ(T)}{\eta(T, RH)} \right)^\xi \quad (S1.2)$$

30 In this equation,  $D_{H_2O}^\circ(T)$  is the self-diffusion of water calculated with the Stokes–Einstein equation at 293 K  
31 ( $2.15 \times 10^{-9} \text{ m}^2 \text{ s}^{-1}$ ).  $\eta_{H_2O}^\circ(T)$  is the viscosity of pure water obtained from literature with values of  $1.78 \times 10^{-3} \text{ Pa}\cdot\text{s}$   
32 at 273 K,  $1.31 \times 10^{-3} \text{ Pa}\cdot\text{s}$  at 283 K,  $1 \times 10^{-3} \text{ Pa}\cdot\text{s}$  at 293K, and  $8.14 \times 10^{-4} \text{ Pa}\cdot\text{s}$  at 303 K. (Weight, 2019).  $\eta(T, RH)$  is  
33 the measured viscosity of sucrose at the corresponding RH and temperature.  $\xi$  is the fractional exponent, using Eq.  
34 (S1.3), which accounts for the relative size of the diffusing molecule and the matrix:

$$\xi = 1 - \left[ A \times \exp \left( -B \frac{r_{diff}}{r_{matrix}} \right) \right] \quad (S1.3)$$

35 where coefficient values of  $A = 0.73$  and  $B = 1.79$  (Evoy et al., 2020),  $r_{diff}$  hydrodynamic radius of water of 0.1  
36 nm. (Price et al., 2016) The matrix radius,  $r_{matrix}$ , of the sucrose molecule was calculated under the assumption of  
37 spherical geometry, using  $342.30 \text{ g mol}^{-1}$  and a density of  $1.67 \text{ g cm}^{-3}$  (Haynes, 2016), resulting in  $r_{matrix} = 0.44 \text{ nm}$   
38 and  $\xi = 0.51$ . Conditioning durations for the other temperatures (273, 283, 293 and 303 K) are summarized in  
39 Table S1.

40 In addition to water, we also estimated the mixing times of organic molecules ( $\tau_{mix,org}$ ) to assess potential diffusion  
41 limitations within the particle matrix itself. This is particularly relevant under conditions where semi-solid or glassy  
42 states may impede equilibration. The diffusion coefficient of organics ( $D_{org}$ ) was calculated using the classical  
43 Stokes–Einstein equation (Eq. S1.4): (Evoy et al., 2019; Evoy et al., 2020)

$$D_{org}(T, RH) = \frac{kT}{6\pi\eta_{org}(RH, T)R_{diff}} \quad (S1.4)$$

44 where  $k$  is the Boltzmann constant,  $\eta_{org}$  is the temperature-and-RH dependent viscosity of OA, and  $R_{diff}$  is the  
45 radius of the diffusing molecules. We used 0.4 nm for  $R_{diff}$ , consistent with typical organic molecules of similar  
46 molecular weight and density. Using the estimated diffusion coefficients, the mixing time of organics was  
47 calculated as:

$$\tau_{mix,org} = \frac{d_p^2}{4\pi^2 D_{org}(T, RH)} \quad (S1.5)$$

48 A particle diameter of 200 nm was used in all calculations, consistent with the typical size of secondary organic  
49 aerosol (SOA) particles observed in the atmosphere (Riipinen et al., 2011; Pöschl et al., 2010). The  $\tau_{mix,org}$  value  
50 represents the time required for the concentration of the diffusing species at the particle’s center to attain 1/e of the  
51 equilibrium concentration.  
52

53 The diffusivity coefficient of  $N_2O_5$  in OA ( $D_{N_2O_5}$ ), which is needed in calculations of  $N_2O_5$  uptake coefficient, is  
54 also determined using the fractional Stokes-Einstein equation (S1.2 and S1.3). In Eq. S1.3, the hydrodynamic radius  
55 of  $N_2O_5$  is set to be 0.25 nm (Grzanic et al., 2015).

## 56 S2. Viscosity parameterization of sucrose-H<sub>2</sub>O droplets

57 We parameterized the viscosity of sucrose-H<sub>2</sub>O droplets as a function of RH and temperature, following the  
58 approach proposed by Kiland et al. (2023) Experimental data at  $293 \pm 1$  K was fitted using a mole-fraction-based  
59 Arrhenius mixing rule (Eq. S2.1)

$$\log(\eta(\text{RH}, 293\text{K}) = x_{\text{org}} \log_{10}(\eta_{\text{org,dry}}) + (1 - x_{\text{org}}) \log(\eta^{\circ}_{\text{H}_2\text{O}}) \quad (\text{S2.1})$$

60 where  $\eta(\text{RH}, 293\text{K})$  is the viscosity of the mixture at room temperature,  $\eta_{\text{org,dry}}$  is the viscosity of the organic  
61 component at 0% RH ( $1 \times 10^{12}$  Pa·s), and  $\eta^{\circ}_{\text{H}_2\text{O}}$  is the viscosity of pure water ( $1 \times 10^{-3}$  Pa·s).

62 The mole fraction of organic components,  $x_{\text{org}}$ , was calculated from the mass fraction using:

$$x_{\text{org}} = \frac{\frac{w_{\text{org}}}{M_{\text{org}}}}{\frac{w_{\text{org}}}{M_{\text{org}}} + \frac{1 - w_{\text{org}}}{M_{\text{H}_2\text{O}}}} \quad (\text{S2.2})$$

63  
64 where  $w_{\text{org}}$  is the mass fraction of the organic component, and  $M_{\text{org}}$  and  $M_{\text{H}_2\text{O}}$  are the molecular weights of sucrose  
65 and water, respectively.  $w_{\text{org}}$  was determined from the water activity ( $a_w = \text{RH}/100$ ) using a mass-based  
66 hygroscopicity parameter  $k$ , as shown in Eq. S2.3.

$$w_{\text{org}} = \left( 1 + k \left( \frac{a_w}{1 - a_w} \right) \right)^{-1} \quad (\text{S2.3})$$

67  
68 To evaluate the tropospheric distribution of predicted sucrose-H<sub>2</sub>O droplet phase state and mixing times, we  
69 utilized monthly mean reanalysis data for temperature and RH from the Copernicus Climate Data Store  
70 (<https://cds.climate.copernicus.eu/>) (Hersbach et al., 2023). The dataset provides global coverage from 90°N to  
71 90°S and 180°W to 180° E, with temporal resolution spanning January 2020 to December 2024. The altitude ( $h$ )  
72 corresponding to each pressure level was computed using the following equation derived from the barometric  
73 formula:

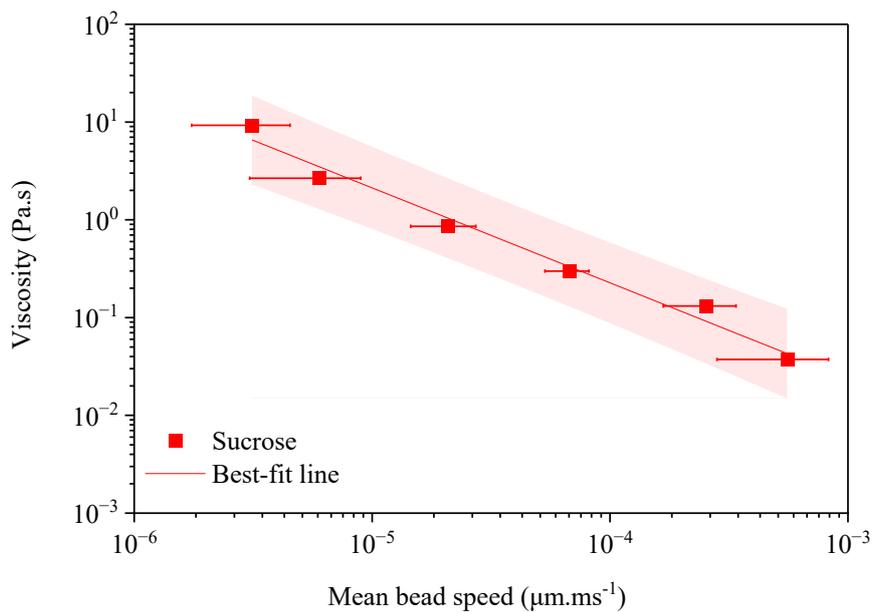
74 Fitting the experimental data at 293 K yielded the value of  $k = 0.061 \pm 0.0023$ . As shown in Figure S8, the model  
75 provides a good fit to the experimental data, supporting the robustness of the parameterization. Peters and  
76 Kreidenweis (2007) reported  $k$  values between 0.01 and 0.5 for diverse organic compounds, highlighting variability  
77 in hygroscopicity with molecular structure.

78 **S3. Collection of global ambient zonal-RH and -temperature**  
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$$h = \frac{1 - \left(\frac{P}{P_o}\right)^{\frac{R\lambda}{M_{air}}}}{\lambda/T_o} \quad (3)$$

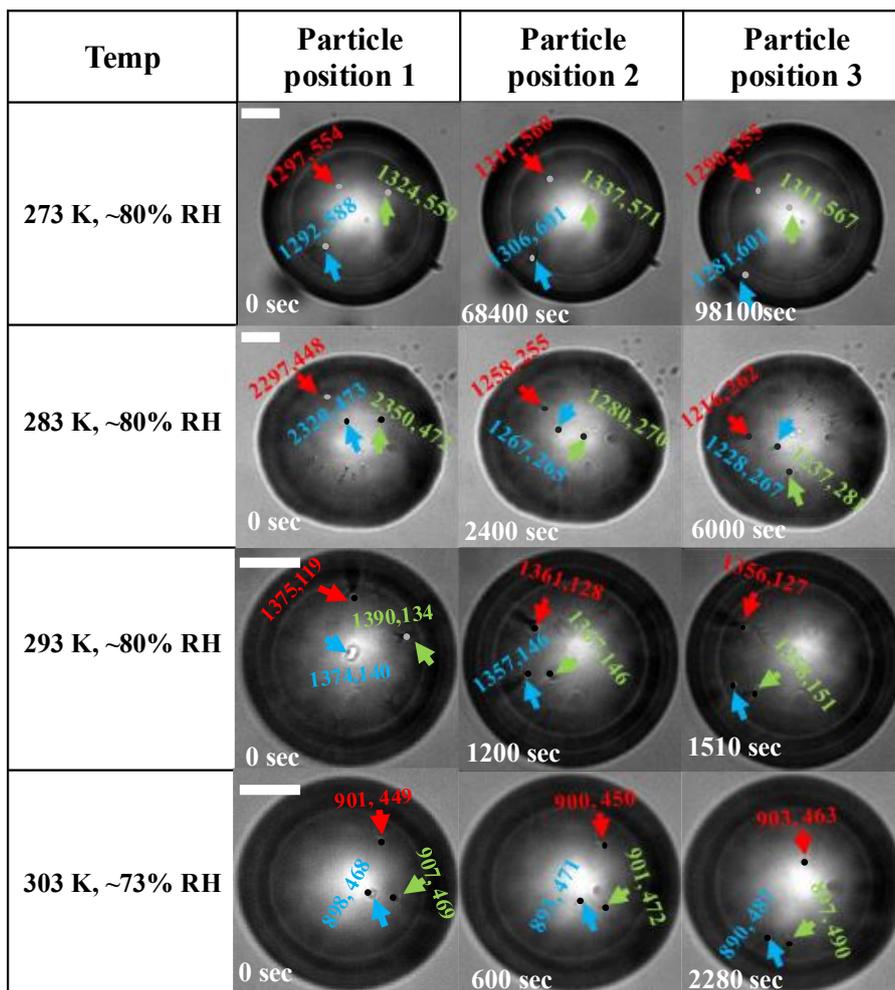
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81 where  $P$  is the pressure at the given level (1000 hPa – 100 hPa),  $p_o$  is the pressure at sea level (101325 Pa),  $R$  is the  
82 gas constant (8.314 J mol<sup>-1</sup> K<sup>-1</sup>),  $\lambda$  is the temperature lapse rate (6.5 K km<sup>-1</sup>),  $M_{air}$  is the molecular mass of the air  
83 (28.97 g mol<sup>-1</sup>), and  $T_o$  is the mean temperature at the surface (288.15 K). To construct latitude–altitude profiles,  
84 the monthly RH and temperature values were averaged across all longitudes at each latitude and pressure level.  
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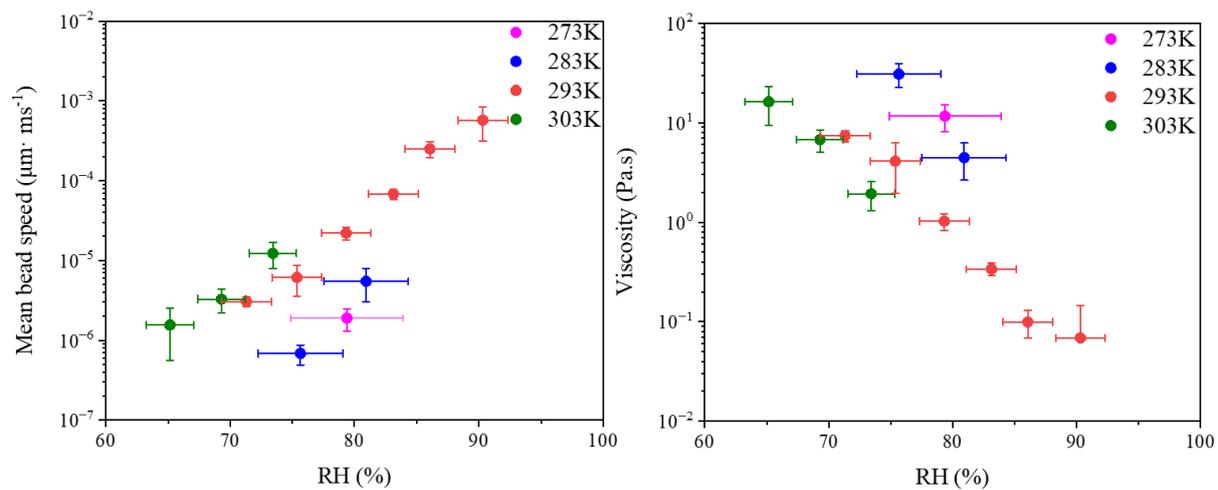
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 88 **Figure S1: Calibration line illustrating the relationship between mean bead speeds and viscosities of sucrose-H<sub>2</sub>O**  
 89 **droplets at varying relative humidity (RH) levels. A linear regression, shown by the red dotted line, fits the data with the**  
 90 **equation:  $viscosity = 0.00003 \times (mean\ bead\ speed)^{-0.971}$ . The pink area denotes 95% prediction bands of fitting**  
 91 **to the data in this study. The uncertainty in mean bead speed along  $x$ -axis is calculated from standardization of 2 – 5**  
 92 **beads within 3 – 5 particles for each RH value.**

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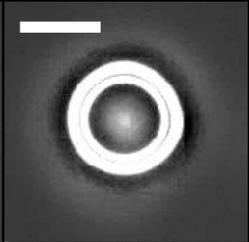
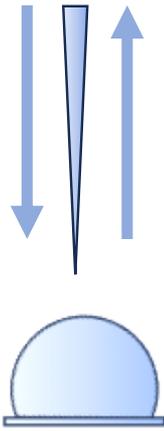
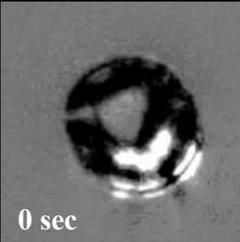
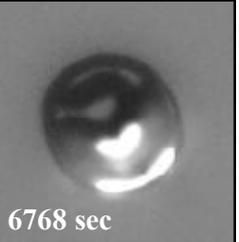
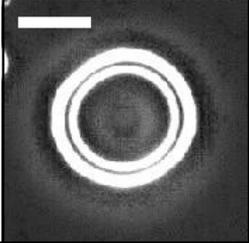
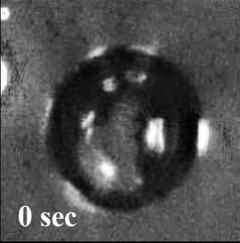
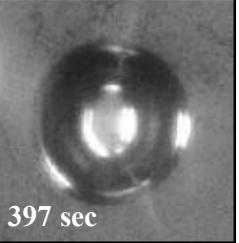
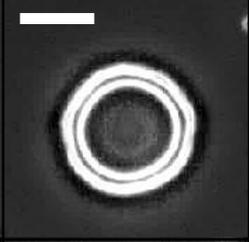
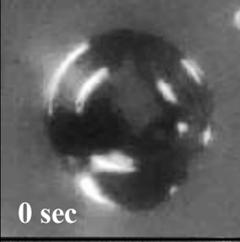
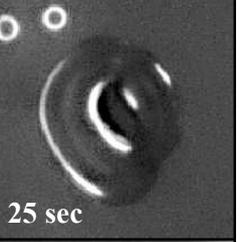
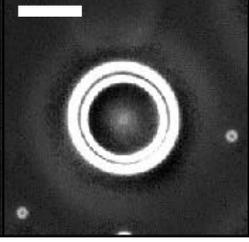
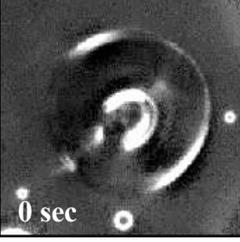
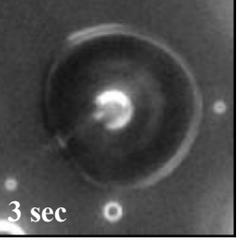
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 95 Figure S2: Optical images of sucrose-H<sub>2</sub>O droplets during a typical bead-mobility experiment at different temperatures.  
 96 Three labeled beads with tracked x and y coordinates were used to determine average bead speeds using ImageJ  
 97 software. The size of the scale bar is 20  $\mu$ m.

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 100 **Figure S3: (a) Mean bead speed and (b) resulted viscosities from bead-mobility experiments for sucrose-H<sub>2</sub>O droplets**  
 101 **as a function of temperature and relative humidity (RH). The x-axis error bars represent the uncertainty in RH from**  
 102 **the RH sensor during calibration at each temperature, while y-axis error bars indicate the standard deviation of the**  
 103 **measured bead speeds and viscosity calculated from 3 – 5 beads across 2 – 5 particles at each RH level.**

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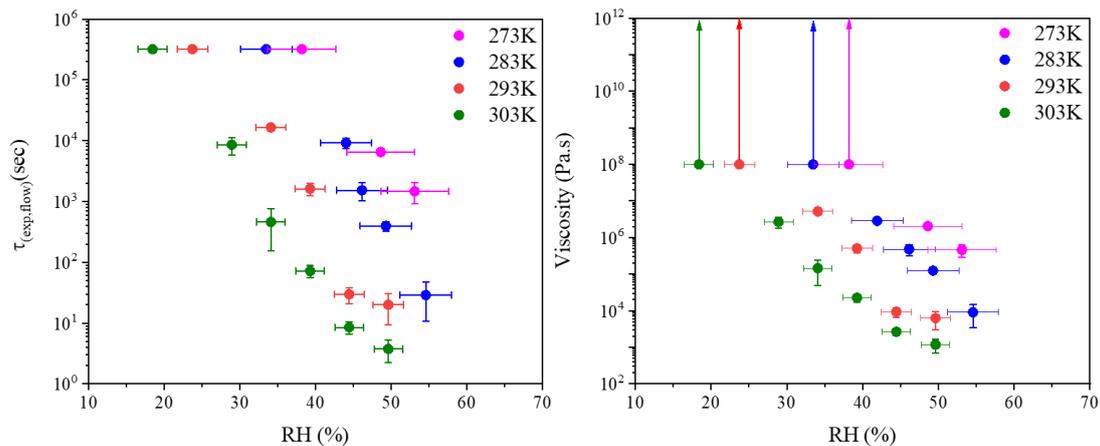
Temp	Pre-poking	Poking	Post-poking	$\tau_{\text{exp,flow}}$
273K, ~ 50% RH			 0 sec	 6768 sec
283 K, ~ 50% RH			 0 sec	 397 sec
293 K, ~ 50% RH			 0 sec	 25 sec
303 K, ~ 50% RH			 0 sec	 3 sec

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Figure S4: Optical images for experimental flow times ( $\tau_{\text{exp, flow}}$ ) during poke-and-flow experiments at different temperatures. White scale bar indicates 20  $\mu\text{m}$ .



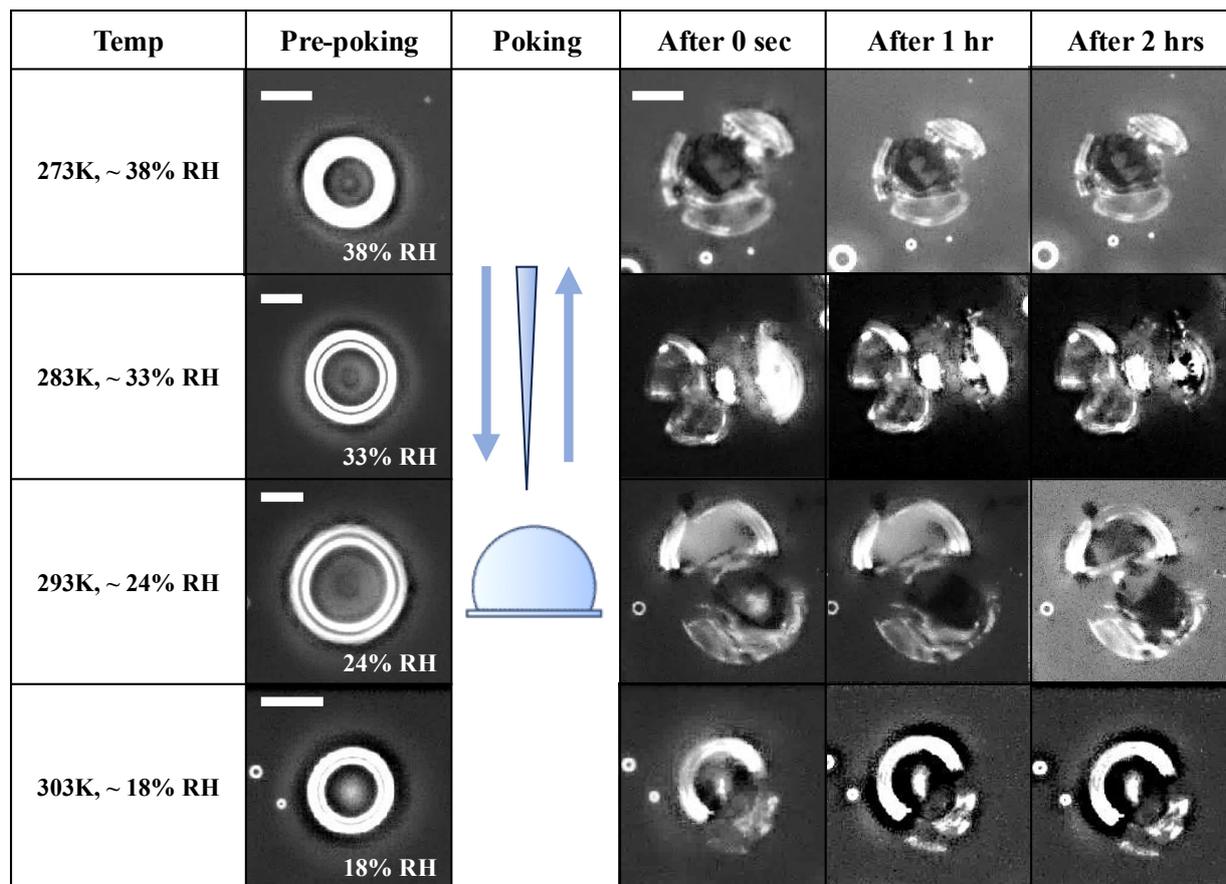
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109 **Figure S5: (a) Mean experimental flow time ( $\tau(\text{exp, flow})$ ) from poke-and-flow experiments for sucrose-H<sub>2</sub>O droplets as**  
 110 **a function of temperature and relative humidity (RH). The x-axis error bars represent the uncertainty in RH from the**  
 111 **RH sensor during calibration at each temperature, while y-error bars indicate the standard deviation of the measured**  
 112  **$\tau(\text{exp, flow})$ , calculated from measurements of 3 – 4 particles at each RH level. (b) Resulted viscosities from the  $\tau(\text{exp, flow})$**   
 113 **and the equation proposed by Sellier et al. (2015) The y-error bars indicate the standard deviation of the measured**  
 114 **viscosity.**

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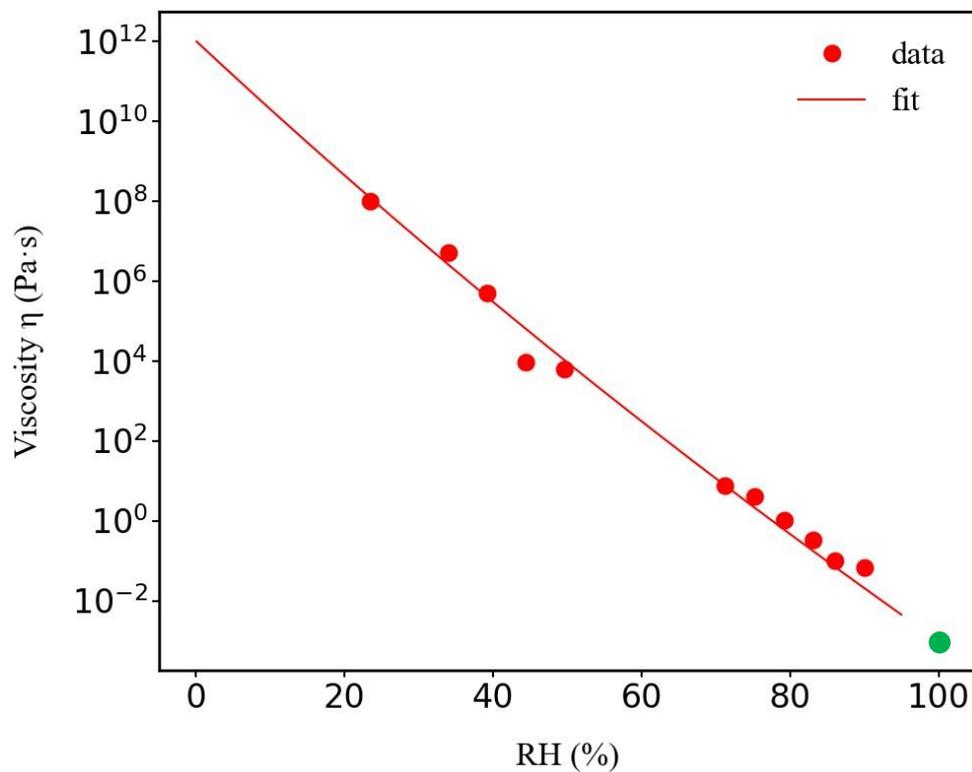
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 119 **Figure S6: Optical images of sucrose-H<sub>2</sub>O droplets during poke-and-flow experiments at different temperatures.**  
 120 **Observations were made when particles cracked at certain relative humidity (RH), and they were then monitored for**  
 121 **longer than 2 hours, with no evidence of flow restoration was detected. RH was consistently regulated throughout the**  
 122 **pre-poking, poking, and post-poking stages. White scale bars represent 20 μm.**

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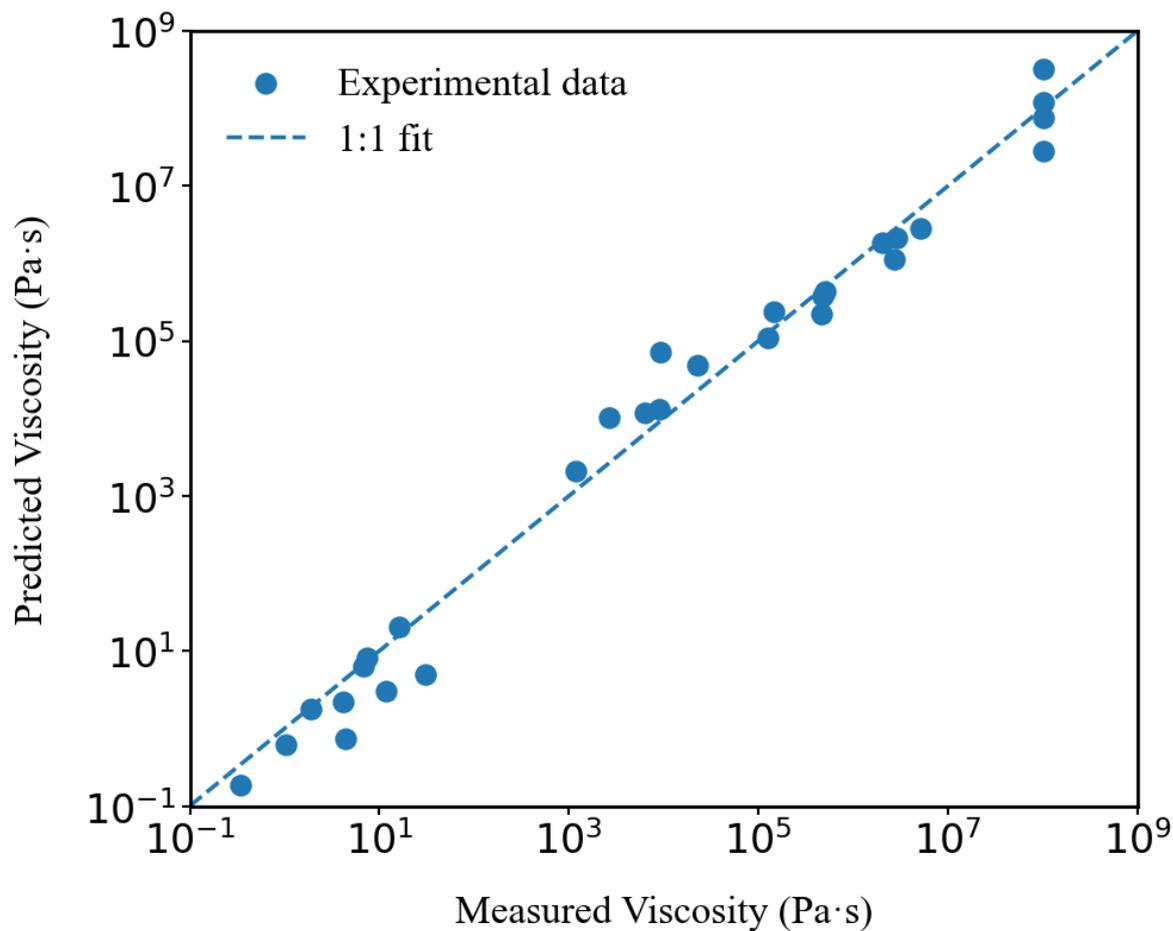
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126 **Figure S7: The log viscosity of sucrose-H<sub>2</sub>O droplets as a function of RH at room temperature. The solid line is a mole-**  
127 **fraction-based Arrhenius mixing rule fit to the viscosity data (eq S2.1), which yields a hygroscopicity parameter,  $\kappa =$**   
128 **0.061. The point at 100% RH (green circle) represents the viscosity of water at room temperature.**

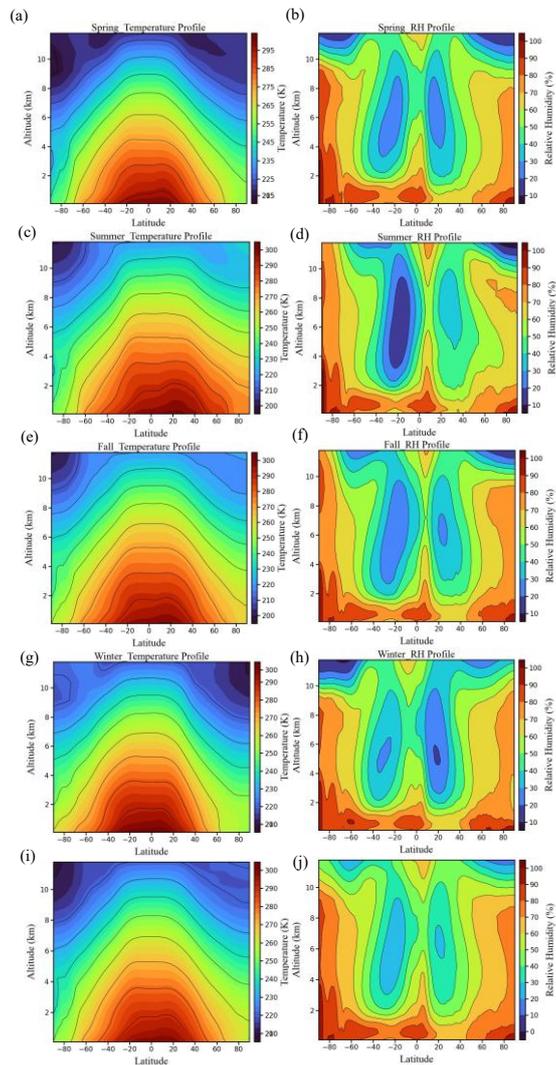
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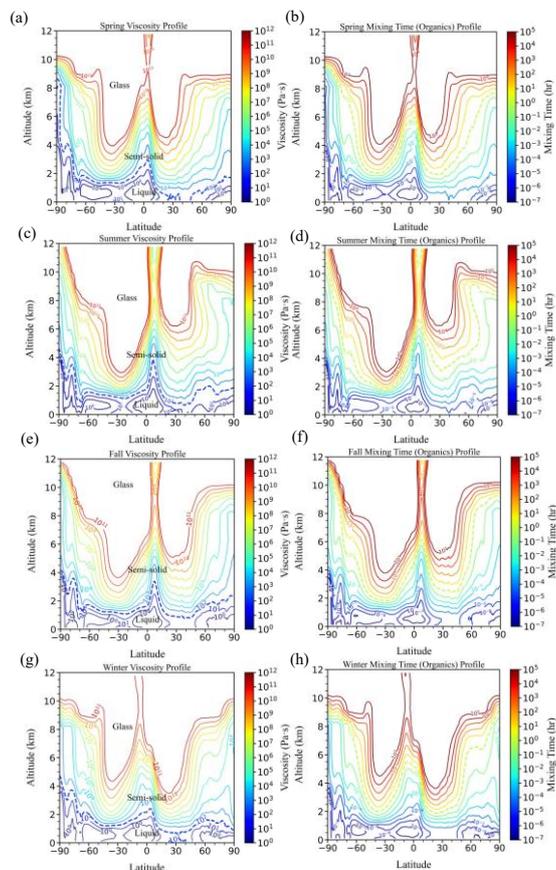
131 **Figure S8: Comparison of experimental and predicted viscosity values for sucrose-H<sub>2</sub>O droplets. The experimental**  
 132 **values (blue circles) were measured between 273 and 303 K and ~20 to ~90 % RH, whereas predicted viscosities (dashed**  
 133 **line) were calculated using the VFT equation. The fitting procedure yields the fragility parameter ( $D_f$ ) value of  $13 \pm 1$ .**

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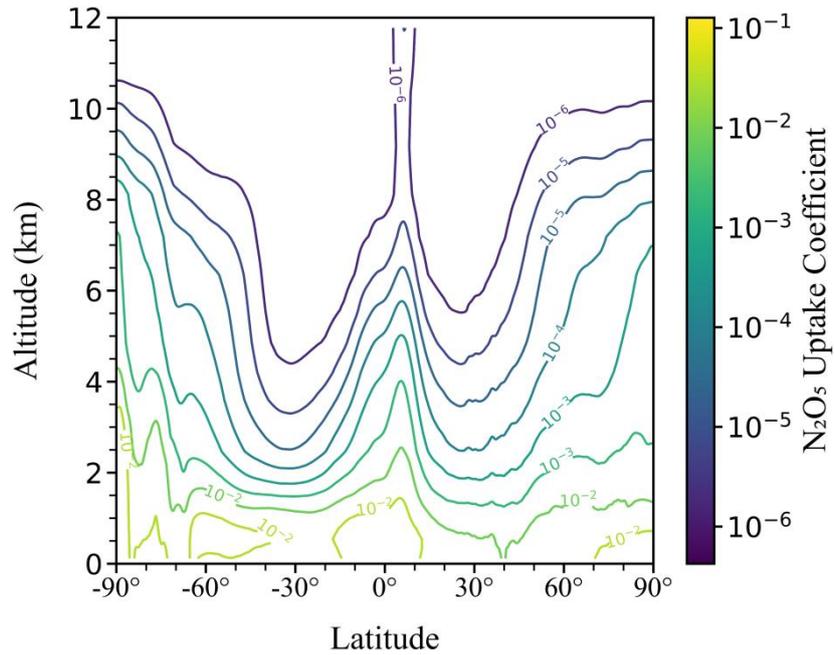
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136 **Figure S9: Altitude–latitude profiles of zonal-mean temperature and RH obtained from Copernicus Climate Data Store**  
 137 **(<https://cds.climate.copernicus.eu/>), averaged seasonally and annually over the period 2020-01-01 to 2024-12-01. Panels**  
 138 **present seasonal profiles for spring (a and b), summer (c and d), fall (e and f), and winter (g and h), showing temperature**  
 139 **and RH, respectively. Panels (I and j) display the annual average zonal-mean profiles of temperature and RH.**



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 141 **Figure S10: Seasonal zonal-mean profiles of viscosity and mixing time of sucrose droplets as a function of altitude and**  
 142 **latitude, derived from temperature and RH data obtained from Copernicus Climate Data Store**  
 143 **<https://cds.climate.copernicus.eu/> and averaged over the period 2020-01-01 to 2024-12-01. Panels (a and b) present**  
 144 **viscosity and mixing time for spring, panels (c and d) for summer, panels (e and f) for fall, and panels (g and h) for**  
 145 **winter. The blue dashed line shows the transition from liquid to semi-solid state. The light green dashed shows the mixing**  
 146 **time of 1 hr.**

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 149 **Figure S11: N<sub>2</sub>O<sub>5</sub> uptake coefficient in 200 nm sucrose droplets as a function of altitude and latitude based on annual**  
 150 **average zonal-mean RH and temperature fields for the years 2020 to 2024, obtained from Copernicus Climate Data**  
 151 **Store (<https://cds.climate.copernicus.eu/>). Surface hydrolysis is not considered in calculations of N<sub>2</sub>O<sub>5</sub> uptake coefficient**  
 152 **( $\Gamma_s$  is set to be 0 in Eq. 3).**

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169 **Table S1.** Summary of the conditioning time for sucrose-H<sub>2</sub>O droplets at different temperature and RH conditions.  
 170  $\tau_{mix,H_2O}$  values are the calculated characteristic mixing time of water within the sucrose-H<sub>2</sub>O droplets based on the  
 171 calculated diffusion coefficient of water.  $\tau_{conditioning}$  values represent the experimental conditioning time.  $\tau_{conditioning}$   
 172 /  $\tau_{mix,H_2O}$  ratio values exceeding 1 indicates that the particles have likely reached equilibrium with the surrounding  
 173 RH.  
 174

RH	$\tau_{conditioning}$ (h)	$\tau_{mix,H_2O}$ (h)	$\tau_{conditioning} / \tau_{mix,H_2O}$
<b>Temp = 273 K</b>			
79	3.30E-01	3.64E-03	90.6
53	1.50E+00	2.39E-01	6.3
49	1.50E+00	4.58E-01	3.3
<b>Temp = 283 K</b>			
81	1.60E-01	1.85E-03	86.5
76	1.60E-01	5.00E-03	32.0
55	1.00E+00	2.19E-02	45.7
49	1.00E+00	8.86E-02	11.3
46	1.00E+00	1.94E-01	5.1
42	1.00E+00	4.64E-01	2.2
<b>Temp = 293 K</b>			
90	1.60E-01	1.84E-04	868.3
86	1.60E-01	2.23E-04	718.4
83	1.60E-01	4.18E-04	383.0
79	1.60E-01	7.37E-04	217.2
75	1.60E-01	1.50E-03	106.4

71	1.60E-01	2.03E-03	78.7
50	5.00E-01	1.82E-02	27.5
44	5.00E-01	2.14E-02	23.3
39	5.00E-01	1.48E-01	3.4
34	5.00E-01	5.08E-01	1.0

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**Temp = 303 K**

73	1.60E-01	8.87E-04	180.5
69	1.60E-01	1.69E-03	94.9
65	1.60E-01	2.65E-03	60.5
50	5.00E-01	6.54E-03	76.5
44	5.00E-01	8.78E-03	56.9
39	5.00E-01	2.69E-02	18.6
34	5.00E-01	5.95E-02	8.4
29	5.00E-01	3.25E-01	1.5

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