

Please see below the point-by-point responses to Referee 1, and the actions taken regarding their concerns.

Major changes

The revised manuscript includes relevant changes with respect to the submitted manuscript in order to address the reviewer's concerns. Please note that, due to the reorganization of certain sections, subsection numbering and figure references have been updated accordingly in the revised manuscript. In addition, we have replaced "predominant" with "dominant" in the title for grammatical correctness. In the text that follows, the suggestions and comments of the Referees are in black and plain font, and our responses are in *italics and blue font*.

Referee 1

General comment:

In this manuscript, the authors report nitrous oxide production pathways in two eutrophic reservoirs using ^{15}N tracer incubations, natural-abundance isotopes, and functional gene quantification. The methodologies are well-established and proven effective in clarifying nitrous oxide production in aquatic systems; the datasets add valuable observations to the community; and the analyses generally support the claim that denitrification dominates water-column nitrous oxide production. The manuscript is suitable for target journal when properly revised. Note: line numbers and pages from authors' PDF file.

My main concern is in the title. The authors argue that denitrification could be the predominant process regulating nitrous oxide production, and in the Discussion section, the authors present evidence about higher nirS abundance and isotopic patterns associated with denitrification. However, this argument may be weakened by the fact that, denitrification was generally absent at the oxygenated surface water (figs. 2 & 3), and the N_2O consumption potential inferred from natural abundance isotopic data. The authors acknowledged that *nosZ* was quantified only at the deepest depths ($n=4$) and therefore cannot constrain N_2O reduction within the entire reservoir (line 270–275; p.12 figure caption). I suggest the authors reframe the argument/statement to better characterize the novelty of their work.

We thank the reviewer for their thoughtful comment. We believe there may be a misunderstanding regarding the treatments applied during the incubation experiments. Specifically, the incubations with $^{15}\text{NO}_3^-$ to determine N_2O production by denitrification were not conducted in the epilimnion; rather, they were performed in the oxycline and hypolimnion. This was described in the Methods section (lines 173-174) and summarized in Table 1.

"The first treatment ($^{15}\text{N}-\text{NH}_4^+ + ^{14}\text{N}-\text{NO}_3^-$) was performed at all the depths ($n=12$), but the second treatment ($^{15}\text{N}-\text{NO}_3^- + ^{14}\text{N}-\text{NH}_4^+$) was performed only at the oxycline and hypolimnion ($n=7$, Table 1)."

To avoid confusion, we have updated the figures to clearly indicate that the treatment with $^{15}\text{NO}_3^-$ used to detect denitrification was not performed in the epilimnion (see Fig. 2 below as an example, which corresponds to Fig. 3 in the revised manuscript). Therefore, we cannot conclude that denitrification was absent in oxygenated surface waters, only that it was not measured there.

*As the reviewer correctly noted, natural abundance data suggest that N_2O may also be consumed by denitrification, and we detected *nosZ* in the analyzed samples. However, despite*

potential N_2O consumption, the overall effect of denitrification remains net production, as evidenced by the substantial accumulation of N_2O in the water column. We are aware that based on those results, we cannot constrain the N_2O consumption in the entire reservoir but can identify the net production trends.

In order to reframe the argument/statement and highlight the novelty of this work, we have done the following changes:

Abstract:

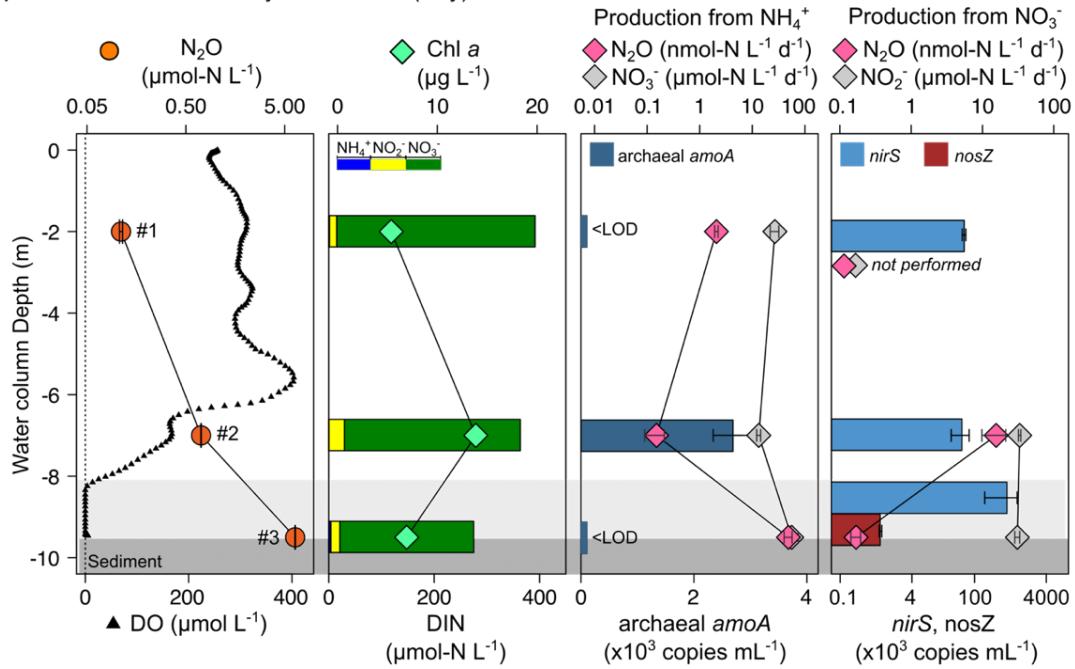
*“Reservoirs are important sites for nitrogen cycling and a significant global source of the potent greenhouse gas nitrous oxide (N_2O). They receive nitrogen inputs from agriculture and urban sources, fueling N_2O production via nitrification, denitrification, and photochemodenitrification. However, existing estimates of N_2O production in reservoirs remain uncertain because most studies have focused on N_2O in rivers or lake sediments, often overlooking the water column of lentic systems. Here, we present the first integrated assessment of N_2O production pathways in reservoir water columns using stable isotope tracer incubations alongside analyses of in situ natural abundance of nitrogen pools and functional genes involved in nitrification (*amoA*) and denitrification (*nirS*), across two eutrophic reservoirs with contrasting morphometries. We used ^{15}N - NH_4^+ and ^{15}N - NO_3^- tracers to quantify rates of N_2O production, nitrification, and nitrate reduction at the beginning and the end of the stratification period. Notably, nitrate concentration decreased by up to 49 % over the two months. N_2O production from ammonium ranged from 0.02 to 48.6 nmol-N $L^{-1} d^{-1}$, while N_2O from nitrate varied from 0.2 to 61.0 nmol-N $L^{-1} d^{-1}$. High rates of nitrification, nitrate reduction to nitrite, and rapid nitrite turnover were observed, with total N_2O production significantly correlated with *nirS* gene abundance. A strong positive correlation was found between $\delta^{15}N$ - NO_2^- and both N_2O concentration and *nirS* abundance. These findings reveal that denitrification and nitrite dynamics play a central role in N_2O formation within reservoir water columns, advancing understanding of nitrogen loss and greenhouse gas emissions from lentic systems.”*

Introduction (Lines 66 – 76 in the revised manuscript):

“Shallow systems tend to emit N_2O continuously due to weak thermal stratification and less capacity to accumulate N_2O . Further studies on N_2O production in the water column of reservoirs with different morphometries are required to improve our knowledge of N_2O emissions. To address this gap, we present the first integrated assessment of N_2O production pathways in reservoir water columns, combining stable isotope tracer incubations with analyses of in situ natural abundances of the N pools and functional genes involved in N_2O cycling to quantify N_2O production rates and trace the origin of the N_2O in the water column of two reservoirs. We used ^{15}N - NH_4^+ to quantify the rates of N_2O production from NH_4^+ , and ammonia oxidation to nitrite and nitrate; and ^{15}N - NO_3^- to trace the formation of N_2O and NO_2^- from NO_3^- reduction. Incubations were performed at three depths at the beginning and end of summer stratification. We selected a shallow and a deep reservoir (Cubillas and Iznájar, respectively) located in watersheds with high N inputs, both of them monomictic with significant emissions and concentrations of N_2O (León-Palmero et al. 2020a, 2023), providing an ideal setting to explore N_2O cycling.”

Figure 2 revised (Figure 3 in the revised manuscript):

(a) Cubillas reservoir in early stratification (July)



(b) Cubillas reservoir in late stratification (September)

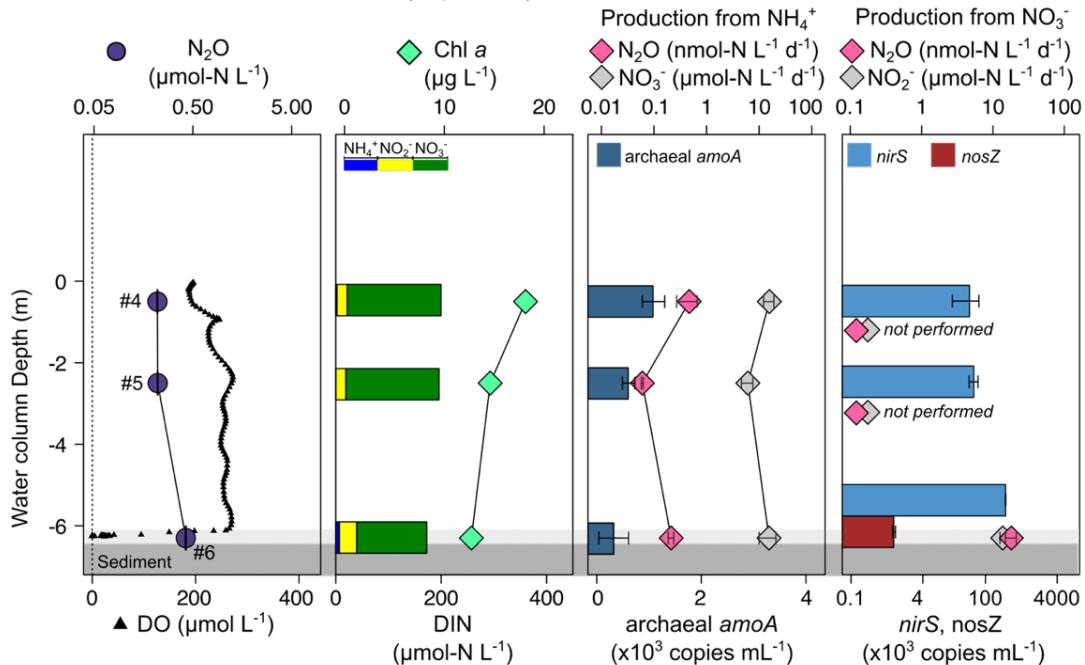


Figure 3. Vertical profiles of the N_2O concentration, production rates, marker genes (colored bars), and other relevant biogeochemical variables in the Cubillas reservoir in July (a) and September (b). Dissolved N_2O ($\mu\text{mol-N L}^{-1}$, mean \pm standard error), and DO concentration ($\mu\text{mol L}^{-1}$); Chl a concentration ($\mu\text{g L}^{-1}$), and DIN concentration ($\mu\text{mol-N L}^{-1}$); N_2O production ($\text{nmol-N L}^{-1} \text{d}^{-1}$) and nitrification (NO_3^- production, $\mu\text{mol-N L}^{-1} \text{d}^{-1}$) from NH_4^+ ; N_2O production ($\text{nmol-N L}^{-1} \text{d}^{-1}$) and NO_2^- production ($\mu\text{mol-N L}^{-1} \text{d}^{-1}$) from NO_3^- , and the abundance of the target genes ($\times 10^3$ copies mL^{-1} , mean \pm standard deviation). Numbers next to N_2O concentrations refer to the sample ID in Table 1. The light gray area represents the suboxic zone (DO $< 10 \mu\text{mol L}^{-1}$) and the dark gray the sediment. <LOD means below level of detection. Note the logarithmic scales for some panels. nosZ gene abundance was only determined in the deepest layers. N_2O and NO_2^- production were only determined in the oxycline and hypolimnion.

Some minor comments below:

Line 31 – 33: These global estimates of increased N₂O emissions from inland waters are often with large uncertainties. Better to state as “mean \pm uncertainty”

We agree with the reviewer, but unfortunately, the original article does not provide such uncertainties. Please see the original article by Wang et al. (2023)

Original text from Wang et al. (2023): “The increase in reservoir emission (0.44 Tg N yr⁻¹) during 1900–2010 accounts for 50% of the total increase in inland-water emission (0.89 Tg N yr⁻¹), indicating that reservoirs are currently an important source of N₂O to the atmosphere”.

Wang, J., Vilmin, L., Mogollón, J. M., Beusen, A. H. W., van Hoek, W. J., Liu, X., Pika, P. A., Middelburg, J. J., and Bouwman, A. F.: Inland Waters Increasingly Produce and Emit Nitrous Oxide, Environ. Sci. Technol., 57, 13506–13519, <https://doi.org/10.1021/acs.est.3c04230>, 2023.

Line 35 – 37: There are increasing number of literatures about GHG emission from reservoirs (doi: 10.5194/bg-11-5245-2014; 10.1016/j.watres.2025.123420; 10.3390/su132111621)

We thank the reviewer for their suggestion. We have included the finding by Chen et al., (2025) in the introduction. In Sturm et al., (2014), no production was detected during sediment incubations. In Ion and Ene (2021) authors focus the GHG assessment on CO₂ and CH₄ emissions, so this paper does not provide useful background information about our subject, N₂O.

Please see the revised text below: “A recent study estimated that N₂O accounted for more than 80% of the total GHG emissions from hydroelectric reservoirs in China in 2020 (Chen et al., 2025).” Lines 41 and 42 in the revised manuscript.

Line 52 – 54: Should clarify about low but not zero oxygen promoting partial denitrification and thus net N₂O production.

We included that information in the introduction (Lines 55 – 60 in the revised manuscript):

“Denitrification is an anaerobic pathway, and oxygen regulates the activity of the denitrifying enzymes, especially the N₂O reductase (Bonin et al., 1989; Zumft, 1997). Therefore, at low but non-zero oxygen concentrations, N₂O reductase might be inhibited, promoting partial denitrification and resulting in net N₂O production. Moreover, many bacteria can denitrify in both oxic and anoxic conditions (Hochstein et al., 1984; Lloyd et al., 1987), and the presence of denitrifying bacteria has been demonstrated in the oxic and anoxic water column of lakes (Junier et al., 2008; Kim et al., 2011; Pajares et al., 2017) and reservoirs (León-Palmero et al., 2023).”

Line 103: Filtrate passing through 0.7 μ m GF/F filter may not be suitable to characterize DOC; 0.45 μ m filter is recommended.

Water was filtered through pre-combusted (450 °C for 3 hours) Whatman GF/F glass-fiber filters with a nominal pore size of 0.7 μ m, which after combustion have pore size equivalent to 0.45 μ m.

Line 110 – 120: Should mention the pore size of filter collecting molecular samples. Because genetic materials were collected from water pre-filtered through 3 μ m, there may be bias toward underestimating particle-attached nitrifiers or denitrifiers (as stated line 370 – 371).

After the pre-filtration step (3 μm), samples were not collected on a second filter but instead concentrated by centrifugation, following the procedure developed by Boström et al. (2004). This approach is described in detail in the Supplementary Material (Extended Methods: DNA extraction, PCR and qPCR assays).

We have explained that in the main material (Lines 149 - 151):

“We pre-filtered water samples through 3 μm pore size filters, and concentrated the samples by centrifugation, then extracted DNA following Boström et al. (2004), and applied PCR and qPCR to assess presence, and abundance of target genes.”

We agree with the reviewer that pre-filtration through a 3 μm pore-size filter may bias toward underestimating particle-attached microorganisms. We now explicitly acknowledge this limitation in the Discussion, stating:

“Additionally, sample water was pre-filtered before DNA extraction (pore size = 3 μm), which may have excluded microbes attached to particles or suspended sediment, potentially including AOA or Comammox groups.” Lines 413 – 415 in the revised manuscript.

Line 122: It is better to specify the criteria for selecting the three depths in the main text. And specify when the N₂O concentration or isotope samples were measured at which facility. Preservation using mercuric chloride is generally not recommended for sample storage longer than 1 year.

We selected three depths representing the epilimnion, oxycline and hypolimnion or bottom waters of each system. This is explained in the Method **section 2.2 Vertical profiles and Biogeochemical characterization**:

“First, we conducted a vertical profile of the water column using a Sea-Bird 19plus CTD profiler, obtaining continuous measurements of temperature (°C), dissolved oxygen (DO, $\mu\text{mol L}^{-1}$), and conductivity ($\mu\text{S cm}^{-1}$) in the reservoirs. Based on the temperature and DO profiles, we sampled three depths representing the epilimnion, oxycline, and hypolimnion or bottom waters. Water was collected at these three depths using a 5-L UWITEC bottle for further analyses and incubation experiments” Lines 96 – 101 in the revised manuscript.

In the revised manuscript, we have included the information required by the reviewer as follows:

a) N₂O concentration, Method section 2.2 Vertical profiles and Biogeochemical characterization (Lines 101-104 in the revised manuscript)

“Samples for dissolved N₂O analysis were taken in 250-mL air-tight Winkler bottles in duplicate, preserved with a solution of HgCl₂ (final concentration 1 mmol L⁻¹) to inhibit biological activity, and sealed with Apiezon® grease to prevent gas exchange. Samples were stored in the dark at a controlled temperature (25 °C) for less than six months until analysis at the University of Cádiz.”

b) Tracer incubations, Method section 2.5 Experimental setup of ¹⁵N tracer incubations (Lines 176 and 177 in the revised manuscript)

“All samples were stored at room temperature in the dark for less than six months and shipped to the laboratory at Princeton University for further analysis.”

c) Isotopes, Method section 2.3 Natural abundance of stable isotopes ($\delta^{15}\text{N}$ and $\delta^{18}\text{O}$) (Lines 121 and 122 in the revised manuscript)

“Two serum bottles per depth were collected without headspace and poisoned with HgCl₂ to analyze the natural isotopic composition ($\delta^{15}\text{N}$) of the ambient pools of N₂O, NO₂⁻, and NO₃⁻.

Samples were maintained in darkness at room temperature for under six months before shipment to Princeton University for analysis.”

Line 124: What was the volume of air headspace during oxic incubation?

The headspace during the oxic incubation was similar to the anoxic one (=3 mL). This detail has been included in the revised manuscript, as follows:

“Once in the lab, samples from oxic water depths (refer to Table 1) were purged uncapped for 2 min to remove excess N₂O, and a 3-mL headspace with ambient air was maintained after being exposed to ambient air for 30 min.” Lines 161 – 163 in the revised manuscript.

Line 152: The equation (1) by Santoro et al. will overestimate the N₂O production from ⁴⁶N₂O signal. The equation proposed by Ji et al., 2018 GBC (doi: 10.1029/2018GB005887) is recommended.

We thank the reviewer for this suggestion. Both equations (Santoro et al. and Ji et al., 2018) have been published, previously used and are valid for estimating N₂O production. We chose the equation by Santoro et al. because it incorporates the additional factor of 1/F, which is essential for the 46 N₂O term. This factor accounts for the probability that two ¹⁵N atoms will pair to form 46N₂O, which is proportional to 1/F². Including this correction ensures a more accurate representation of the isotopic pairing process.

Similarly, other studies have applied the 1/F² correction to account for the probability of forming doubly labeled molecules, such as ³⁰N₂ in isotope pairing techniques (e.g., Thamdrup and Dalsgaard, 2002). This adjustment ensures accurate calculation of production rates because the likelihood of two ¹⁵N atoms pairing is proportional to F².

Line 186: It seems the incubation timepoints for nitrification were different from those of N₂O production; which two time points were analyzed for ¹⁵N-nitrate?

We used the four time points to study N₂O production from ¹⁵N-Ammonia and from ¹⁵N-Nitrate (t₀, t₁, t₂, t₃), but we used only the first two time points to measure the rates of nitrification and nitrate reduction to nitrite. We have clarified this in the method section:

a) Method section, 2.6 ¹⁵N-N₂O production rates from ¹⁵NH₄⁺ and ¹⁵NO₃⁻ (Lines 184 and 185 in the revised manuscript)

“N₂O production rates for each treatment were calculated from the slope of the increase in mass 45 and 46 during the linear phase over the four time points.”

b) Method section, 2.7 ¹⁵N-NO₂⁻ production & 2.8 ¹⁵N-NO₃⁻ production (Lines 202-203, and 217-218 in the revised manuscript).

“Each rate was calculated from the first two time points, and two or three replicates per time point.”

Line 183 – 190: This section 2.8 about calculating N₂O yields using two equations, (4) & (5); which one is used to represent data described in line 240 & 246?

We used Equation 4 to calculate the yield reported in Line 240 because it refers to the yield during ammonia oxidation. In contrast, the yields presented in Line 245 were calculated using Equation 5, as they correspond to the yield during nitrification. To make this information clearer

to the reader, we have renamed the yields in the method section and in the main text as N_2O -yield_{Amox}, N_2O -yield_{Nit}, and N_2O -yield_{Denit}.

Line 240 (285 in the revised manuscript): “Ammonia oxidation rates (i.e., NO_2^- production from NH_4^+) were only significant in Iznájar’s hypolimnion in September, reaching $215.8 \pm 38.0 \text{ nmol-N L}^{-1} \text{ d}^{-1}$ (N_2O -yield_{Amox}=0.041 %) (Table S2).”

Line 245 (Line 292 in the revised manuscript): “The N_2O yields during nitrification (N_2O -yield_{Nit}) varied from 0.000 to 0.086 %, with the maximum yield observed in the bottom waters of Cubillas in July (Table S2).”

Line 191 – 202: Authors should clarify the consistent amount of N injected into mass spectrometry to determine natural abundance. This is important because the concentrations of N species varied with depth. If varying amounts of N were injected, even for the same water sample, mass spectrometry will yield varying isotopic values.

We appreciate the reviewer’s comment. To clarify, we used the same sample volume for all natural abundance measurements (a 60 mL serum vial after 3-mL of headspace was created).

- a) To measure the natural isotopic composition of the N_2O , we did not adjust the amount of N injected into mass spectrometry but corrected for the mass effect by including standards with a known amount of N_2O gas and internal standards for ^{15}N - N_2O , as described in the method section. Both $\delta^{15}N$ - N_2O (‰) vs. Air- N_2 and $\delta^{18}O$ - N_2O (‰) vs. Vienna Standard Mean Ocean Water (VSMOW) were determined. Isotope measurements were linearity and offset corrected using an internal N_2O reference gas with known isotopic composition (this information was detailed in the methods). Please see subsection [2.3 Natural abundance of stable isotopes \(\$\delta^{15}N\$ and \$\delta^{18}O\$ \)](#) in line 119 in the revised manuscript.
- b) To measure the natural isotopic composition of the NO_2^- , and NO_3^- pools (i.e., $\delta^{15}N$ - NO_2^- and $\delta^{15}N$ - NO_3^-) we converted those compounds to N_2O . NO_2^- was converted to N_2O by using the azide method (McIlvin and Altabet, 2005). We used the denitrifier method to convert NO_3^- to N_2O (Sigman et al., 2001; Granger and Sigman, 2009; Weigand et al., 2016). Both methods and corrections are described in the Method section ([2.3 Natural abundance of stable isotopes \(\$\delta^{15}N\$ and \$\delta^{18}O\$ \)](#), line 119 in the revised manuscript). During the azide method, the sample size was adjusted to contain 10 nmol of NO_2^- , and 20 nmol nitrate NO_3^- . We included known isotope international standards (USGS34 and IAEA N3).

Line 212: Apparently high DO at the surface (> 400 micromolar, ~2-fold saturation) suggests a very strong oxygen source that can only be supported by blooming algal activity. This was not evident from the chl-a concentration profiles (10 – 20 microgram per liter)

Thank you for catching that plotting error, which we have corrected in the revised manuscript. The value in the surface was about $250 \mu\text{mol O}_2 \text{ L}^{-1}$, which is about 100% saturation, but peaked at $400 \mu\text{mol O}_2 \text{ L}^{-1}$ at 5.6 m (Cubillas in July). Line 242 in the revised manuscript.

Line 318–323: The main text should clarify “sample #12 excluded as specified in the caption of Fig. 6f” (as stated in caption L325–333).

We have specified that sample #12 was excluded of the analysis in the revised main text. Please see line 352 in the revised manuscript.

Line 335 & 466: I am not sure about the statement “remove significant amount of N”. There could be possibility about the loss of nitrate due to algal assimilation, and the biomass being exported to the

sediment. In addition, statements about **reservoir-level DIN loss** described in Section 4.4 and the supplementary lack uncertainty estimates and assumptions (depth weighting, temporal representativeness, varying hydraulic retention time, etc). Please add a short paragraph enumerating assumptions and explain the possible caveats. Readers may realize that these DIN loss potential could be specific to the sampling period (Jul–Sep), not necessarily annual rates (L465–474 & Table 2).

Thank you for this observation. We agree that nitrate loss in the water column can occur through multiple pathways, including algal assimilation and subsequent sedimentation of organic matter. Our statement “remove significant amount of N” was intended to refer broadly to fixed nitrogen removal from the water column, not exclusively to denitrification. To clarify, we have revised the text to acknowledge these alternative mechanisms. The revised sentence now reads:

“NO₃⁻ concentration decreased by 49 % and 12 % in Cubillas and Iznájar, respectively, in just two months, which represents a substantial net N loss. This net loss in the water column likely reflects a combination of processes, including denitrification, algal assimilation followed by sedimentation of organic matter, and other biogeochemical transformations.” Lines 371 -374.

We agree with the reviewer that the assumptions made during these calculations should be better stated. Therefore, we included the following paragraph in the section **Extended Methods: Scaling up to the reservoir level**

“The following calculations rely on several assumptions. First, we assume that the July and September profiles are representative of the entire reservoir water column during the stratification period and that depth-weighted concentrations capture vertical variability. Second, we use average reservoir volume to approximate water outflow, without accounting for short-term fluctuations in hydraulic retention time or drawdown dynamics. Third, we assume minimal nitrogen inputs from the watershed during the study period, since summer is the dry period and observed reservoir drawdown (and thus minimal inputs from streams, rain or runoff) but cannot exclude minor contributions from groundwater or episodic events. Finally, these estimates reflect net changes over the sampling interval (July–September) and should not be interpreted as annual rates. Spatial heterogeneity and biological processes such as algal assimilation and sedimentation may also influence apparent DIN loss. Therefore, these values represent potential nitrogen removal only during the stratified season, not whole-year budgets.”

Additionally, we also included the following paragraph in the revised discussion (Lines 525 – 530):

“These estimates represent a major seasonal N loss event rather than annual rates. They are based on DIN concentration differences between July and September, without considering whether the reservoirs received N inputs from their watersheds during that period. Since summer is the dry period, and drawdown of the reservoirs exceeded any input via rain or runoff, N inputs from the watersheds were likely minimal during the study period. Further details on the calculations and assumptions are provided in the Supplementary Material (Extended Methods: Scaling up to the reservoir level).”

Line 362 – 365: Authors should clarify the inconsistencies about anoxic depths with measurable amoA gene abundance (fig 3b), or ammonium oxidation rates (fig 2a).

We have discussed these apparent inconsistencies in the revised discussion (Lines 393 to 399 in the revised manuscript). Please see lines 388 – 404 in the revised manuscript:

“N₂O production rates were significantly correlated with the availability of NH₄⁺ and with nitrification rates, but not with archaeal amoA gene abundance. Despite the hypolimnion of Iznájar in September (#12) being apparently anoxic, we detected a significant production of N₂O from NH₄⁺, ammonia oxidation, and the presence of archaeal amoA genes. This combination of processes and gene detection suggests that trace amounts of oxygen may have been present at levels below the detection limit of our oxygen sensor. Similarly, the presence of trace levels of oxygen may explain the production of N₂O from NH₄⁺, and the nitrification rates in the anoxic waters of Cubillas, although in that case we did not detect the presence of archaeal amoA genes”

Line 374 – 375: Where are the data for ammonia oxidation/nitrification, that can infer comammox?

Initially, we hypothesized that comammox may be relevant in these reservoirs because we detected significant rates of nitrification (complete oxidation of ammonia to nitrate) but we rarely detect any significant rate of ammonia oxidation (ammonia to nitrite). These results are described in the subsection 3.2 Distribution of N₂O production and nitrification rates from ¹⁵N-NH₄⁺. Then, comammox amoA genes were targeted in PCR assays using degenerate PCR primers for clades A and B (Pjevac et al., 2017), but we did not detect them. During these PCR assays, no positive control could be used in these assays, so we cannot completely rule out the presence of these bacteria. This is discussed in the revised text (418 – 423):

“The detection of high nitrification rates, but no significant ammonia oxidation, might suggest that comammox is occurring at these depths. However, our PCR analysis showed no evidence of the presence of comammox bacteria (Fig. S2), although because no positive control was available, we cannot completely exclude their presence. Therefore, we consider the possibility that complete ammonia oxidation could contribute to the observed nitrification rates. Alternatively, we hypothesize that the NO₂⁻ production by ammonia oxidation was tightly coupled to NO₂⁻ consumption by NO₂⁻ oxidizers, such that it could not be detected in the NO₂⁻ pool.”

Line 444: Are the authors suggesting denitrification potential in anoxic water column or the sediment? Please clarify.

In that context, we suggest that denitrification in the water column may be enhanced by the export of organic matter, however, we acknowledge that such export can also stimulate denitrification in the sediment. Although our study focuses on processes occurring in the water column, we recognize that denitrification and/or fixed nitrogen loss in the sediments cannot be ruled out.

In the revised manuscript (lines 492 – 494), we specified that we referred to the water column: “This is also consistent with our previous study in twelve reservoirs (León-Palmero et al., 2023), and may indicate that denitrification is enhanced by particulate material derived from the phytoplankton community in the water column.”

Line 452 – 465: It is better to address the exact profiles shown in figures 2 & 3 when discussing oxycline turnover times and hypolimnetic storage, to help readers understand the statements.

We thank the reviewer for the suggestion. We included the references to the figures in the revised manuscript. Please note that Figs. 2 and 3 became 3 and 4 in the revised manuscript (Lines 502 – 504): “The highest total N₂O production in Cubillas coincided with the highest N₂O concentration at the deepest depth in both months (Fig. 3). In the deeper reservoir, Iznájar, the

highest production was measured at the oxycline, where there is a strong potential for N₂O fluxes, while the highest N₂O concentrations were detected in the hypolimnion (Fig. 4)."

References:

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