



BiasCast: Learning and adjusting real time biases from meteorological forecasts to enhance runoff predictions

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Abstract. The use of deep learning models in hydrology is becoming an ever more prevalent application in operational flood forecasting. Such operational systems face performance degradation when transitioning from high quality reanalysis to meteorological forecast data with lower accuracy. This study investigates training strategies and Long Short-Term Memory network architectures to mitigate forecast-induced bias in maximum daily discharge predictions using the Extended LamaH-CE dataset and a subset of 451 basins. We systematically evaluated cross-domain generalization, transfer learning approaches, Encoder–Decoder LSTMs, Sequential Forecast LSTMs, and the role of input embeddings and integrating past discharge observations. The results show that domain shifts between reanalysis and forecast data lead to substantial skill loss, with median Nash–Sutcliffe Efficiency decreasing from 0.58 to 0.33. Among the tested strategies, the Sequential Forecast LSTM demonstrated the most stable improvements, achieving a median NSE of 0.63. Integrating recent discharge observations further enhanced performance, raising median NSE to 0.71 and surpassing even the reanalysis-driven baseline. In contrast, integrating archived forecasts or using more complex input embeddings did not yield consistent benefits and in some cases degraded model stability. These findings highlight the value of training strategies that allow models to directly learn bias correction during forecast transitions and emphasize the operational potential of combining sequential processing with near real-time discharge observations.

1 Introduction

Accurate runoff prediction stands as one of the most critical challenges in modern hydrology, with far-reaching implications for flood risk management, water resource planning, and the design of resilient hydraulic infrastructure (Beven, 2012; Guo et al., 2021; Tran et al., 2025). While recent advances in deep learning have demonstrated that Long Short-Term Memory Networks (LSTMs) can effectively integrate multiple meteorological datasets to improve runoff simulation accuracy by learning complex spatial and temporal patterns (Kratzert et al., 2021), a fundamental challenge remains: operational forecasting systems rely on biased meteorological forecasts rather than reanalysis or observational data. This dependency introduces a cascade of uncertainties, as meteorological forecasts inherently exhibit lower accuracy and higher uncertainty than observational or reanalysis datasets (Lavers et al., 2021), with forecast errors further amplifying as lead time increases (Nester et al., 2012). The consequences of these uncertainties are particularly severe in flood forecasting applications, where timely and magnitudinal correct runoff predictions are critical for early warning systems and risk management (Chen et al., 2016).

The biases in meteorological forecasts stem from factors such as model resolution, data assimilation techniques, or orographic effects, and they differ depending on the numerical weather prediction model (e.g., ECMWF-HRES, DWD-ICON, NOAA-GFS), the predicted variable itself and the region in question (Haiden et al., 2024). These inaccuracies can propagate through hydrological models and lead to unreliable runoff forecasts, particularly under extreme conditions (Nester et al., 2012). To mitigate this issue, a variety of statistical and machine learning-based bias correction methods have been developed to adjust forecasted meteorological variables prior they are used as input in a hydrological model.



A simple approach to reduce biases in precipitation is described by Lenderik et al. (2007), who are scaling precipitation linearly based on a constant factor calculated from long term observations. To support operational warning systems, Hess (2020) developed the Ensemble Model Output Statistics (Ensemble-MOS) system, which postprocesses ensemble forecasts from COSMO-D2-EPS and ECMWF-ENS. The approach relies on logistic regression and stepwise multiple regression to reduce conditional biases and produce calibrated probabilistic forecasts efficiently. Ko et al. (2020) used the XGBoost machine learning algorithm to correct precipitation forecasts. Their method demonstrates that machine learning can improve rainfall forecasting performance, especially localized heavy rainfall events, which are of special importance for flash floods in small catchments. Zhang et al. (2020) used LSTMs to learn relationships between meteorological forecasts and observed rainfall data. Their results indicate that LSTMs are capable of learning dynamic biases to correct the forecasts from numerical weather predictions and increase forecast reliability, especially for heavy rainfall events. Han et al. (2021) proposed CU-net, a convolutional neural network architecture specifically designed to address systematic biases in gridded numerical weather predictions from ECMWF-IFS. Their grid-based approach represents a methodological advancement by directly correcting spatial forecast fields, enabling comprehensive bias mitigation across continuous meteorological domains. However, the focus on ECMWF-IFS data raises important questions about the correction model's transferability to other numerical weather prediction systems, potentially limiting the generalizability of their bias correction framework and highlighting the need for more robust approaches.

The studies mentioned have in common that the meteorological forecasts are compared either with meteorological station- or reanalysis data. In this context, it is important to note that especially precipitation measurements, whether from rain gauges, radar, or satellite sources, are inherently subject to various sources of uncertainty (Bárdossy et al., 2022). These errors stem from undercatch due to wind effects or sensor limitations (Yang et al., 1999). As a consequence, it can be assumed that even when inputting bias corrected precipitation forecast data to a hydrological model, a source of uncertainty with potential error propagation also arises here, which in turn creates a bias in runoff prediction. In contrast, discharge observations are typically regarded as more reliable compared to precipitation observations, as they represent an integrated hydrological response over the entire catchment and are measured continuously at fixed gauging stations (Hernegger et al., 2015; Mao et al., 2019). Although discharge measurements also carry uncertainty, particularly related to the use of rating curves or sensor malfunction during extreme events, they are less affected by localized measurement errors (De Oliveira and Vrugt, 2022; Villarini et al., 2008).

A method directly improving streamflow forecasts from the physically based Global Flood Awareness System (GloFAS) was developed by Hunt et al. (2022). GloFAS is an operational hydrological forecasting system that couples ECMWF ensemble weather predictions with the LISFLOOD hydrological model to provide streamflow forecasts for rivers worldwide (Alfieri et al., 2013). Instead of bias-correcting the meteorological input variables, Hunt et al. (2022) addressed systematic biases in streamflow forecasts using a statistical bias correction method based on quantile mapping (QM) with spatial optimisation and subsequently applied a damping factor to blend the corrected forecasts with the original raw output. Despite the demonstrated improvements in forecast skill, this bias correction approach has several limitations. First, the quantile mapping correction is dependent on GloFAS forecasts, meaning it is not applicable for regions where no GloFAS forecast is available. Second, the method is lead-time independent, meaning it does not account for the evolution of forecast bias over longer lead times, which can reduce its effectiveness for medium- to long-range forecasts. Third, the applied damping factor, while effective in reducing over-correction, is empirically tuned, which may limit its robustness when applied across diverse catchments or under changing climate conditions. A further limitation of the study is the relatively small number of catchments used (10 gauges), which constrains the generalizability of the findings.

Building on the idea that runoff observations may be more accurate than those of meteorology, Kirchner (2009) proposed a paradigm shift through the concept of "doing hydrology backward," where discharge is used as the primary constraint to infer



the dynamics and uncertainties of upstream processes, such as precipitation or evapotranspiration. Rather than relying solely on uncertain meteorological inputs to predict runoff, backward hydrology extracts information about catchment dynamics directly from the discharge time series itself (Hermegger et al., 2015; Kirchner, 2009). In this respect, the approach could also be used to perform a dynamic bias correction of multiple meteorological forecast variables since runoff data may serve as a more robust target variable in data-driven modelling frameworks than uncertain meteorological observations (e.g. rainfall). Given the hypothesis that large-scale hydrological datasets contain more information than could be described using theoretical or conceptual approaches (Nearing et al., 2021), a way to harness the potential of machine learning is to combine large sample datasets with meteorological forecasts as inputs. In such a setup, the model can learn to assign weights to the forecasts and internally correct their biases, thereby improving the overall runoff prediction accuracy.

In this study we investigate multiple Long Short-Term Memory (LSTM) network architectures and training strategies to reduce meteorological forecast-induced bias in 24-hour ahead maximum daily discharge predictions. The focus on daily maxima ensures that critical peak flows relevant to flood forecasting are not masked by temporal averaging. 24-hour lead time was selected as an initial proof-of-concept to establish baseline performance of bias correction capabilities, as forecast uncertainty generally increases with lead time (Nester et al., 2012), making shorter horizons an appropriate starting point for validating the approach while providing a foundation for future extension to multi-day predictions. We evaluate baseline LSTM configurations, transfer learning approaches, encoder-decoder architectures, and sequential LSTM networks across 451 catchments from the Extended LamaH-CE dataset in Central Europe. Our experiments examine the effectiveness of different data integration scenarios, including the incorporation of past discharge observations and archived forecasts, with the goal of developing robust neural network-based approaches for operational flood forecasting systems that can effectively compensate for systematic biases inherent in numerical weather prediction models.

2 Data and Methods

2.1 Data

This study uses an extended version of the daily LARge-SaMPle DATA for Hydrology and Environmental Sciences for Central Europe (LamaH-CE; Klingler et al., 2021). LamaH consists of 859 gauged catchments including 21 catchment averaged meteorological variables, with more than 60 static catchment attributes. Since the original version of LamaH only contains meteorological ERA5-Land data and Kratzert et al. (2021) show that leveraging multiple meteorological data sources is beneficial in large sample hydrology, we expanded the data by 15 further variables from five sources. The products used are (i) ERA5-Land (Muñoz-Sabater et al., 2021) as in the original LamaH data, (ii) ECMWF-HRES European Center for Medium Range Weather Forecast - High Resolution Forecast (ECMWF, 2025), (iii) E-OBS gridded observational data (Comes et al., 2018), (iv) MSWEP multi-source weighted ensemble precipitation (Beck et al., 2019) and (v) GLEAM global land evaporation Amsterdam model (Miralles et al., 2011). Details of the variables used, including their definitions, units, and sources, are summarized in Appendix A. The data products were obtained as raster data and subsequently aggregated to the LamaH basins. All variables are daily averages (e.g. temperature) or daily sums (e.g. precipitation). For the ECMWF-HRES variables temperature, dew point and sea level pressure, 3 hourly forecast values (8 per day) were calculated as daily averages starting from 0 o'clock (UTC) issue time. A second adaption we made to the LamaH dataset concerns the gauge files. In the daily version of LamaH-CE, there are only the mean daily discharges - we have extracted the daily minima and maxima from the hourly LamaH data for all gauges and extended the daily version with those.

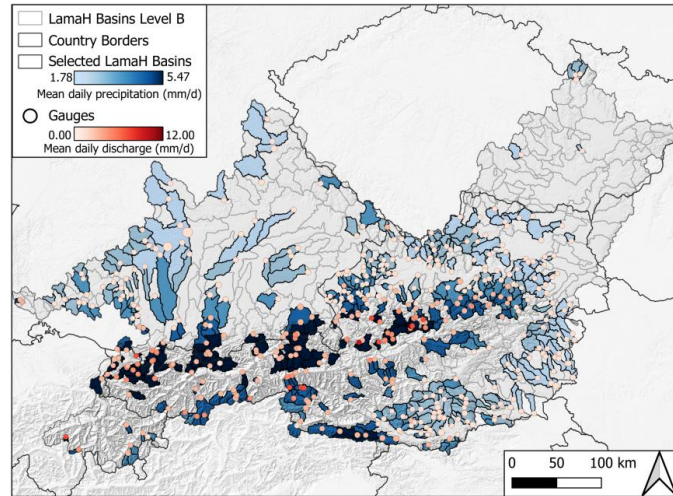


Figure 1. LamaH domain with the 451 subset basins. For better illustration, LamaH subbasins (level B, blue polygons) are shown here, but calculations were performed at LamaH level A (lumped for each gauge). The red points show the runoff gauges located at the catchment outlet.

For the conducted experiments, we used a subset of 451 basins with no and low anthropogenic influence at LamaH aggregation level A, which represents the lumped topographic catchment area of a gauge. Level A is comparable to the aggregation of the catchment areas in the CAMELS (Newman et al., 2015) dataset. The catchments are spatially distributed across the entire LamaH domain, with catchment areas including high alpine-, alpine foothill- and lowland areas.

2.2 Experimental design

To comprehensively evaluate the performance of LSTM-based flood prediction models under different data availability scenarios and training strategies, we designed five distinct experimental groups with the primary research question: **How to reduce the meteorological forecast induced bias in runoff predictions?**

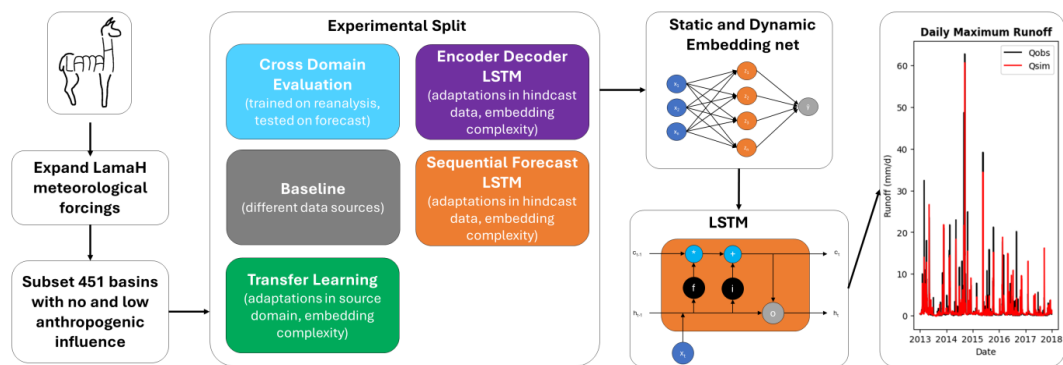


Figure 2. Workflow of the conducted experiments. The LamaH-CE dataset was extended by forecast and further reanalysis data and subset to 451 basins with no and low anthropogenic influence. The experimental split is divided into the deep learning architectures used, followed by a schematic representation of input embeddings for static and dynamic variables which feature space is fed to the LSTM. The last step is the forecast of the daily maximum runoff at the gauges.

All experiments were, in terms of reproducibility, conducted with the NeuralHydrology (Kratzert et al., 2022) python library and trained on different LSTM architectures to predict maximum daily discharge ($q_{\max}$). The models incorporated dynamic meteorological inputs using a 365-day input sequence length, and static catchment attributes (33 physiographic, climatic, and



land cover characteristics), both processed through separate embedding networks. The embedding networks are fully connected neural network layers that transform raw input variables into learned representations, enabling the LSTM to capture non-linear relationships between inputs and allowing different feature combinations to be learned during training rather than being predefined (Ahmed et al., 2023). The experimental framework utilized a consistent temporal split with training data from 2003-2009, validation from 2010-2013, and testing from 2014-2017. Model performance was evaluated using basin averaged Nash-Sutcliffe Efficiency (NSE*, Kratzert et al., 2019b) as the primary loss function. A description of the loss function is attached in Appendix C. Model hyperparameters, such as the number of hidden units, were optimized using Bayesian optimization (see Snoek et al., 2012) with NSE* as the objective function. A detailed description of the performed hyperparameter tuning is attached in Appendix D. Hereafter, we use "domain" equivalent to distinct forcing datasets with unique statistical properties in hydrological modelling to denote a specific data distribution characterized by its feature space and probability distribution in its machine learning sense. Throughout the remainder of this paper, domain specifications and equations reference only dynamic meteorological forcings for clarity, with the understanding that static catchment attributes remain unchanged across all experimental setups.

2.2.1 Baseline

Three baseline experiments were conducted to establish performance benchmarks using different meteorological data sources in a standard LSTM runoff simulation framework. LSTMs are a special form of recurrent neural networks, mainly used for sequential (time series) data (Hochreiter and Schmidhuber, 1997). For a detailed description of the LSTM in relation to hydrological modelling, we refer to Kratzert et al. (2018, 2019a, b). The core tensor equations of the LSTM model responsible for the information flow are presented in Appendix E.

The baseline experiments were conducted using either forecasting data only (FC), reanalysis data only (RA), or a combination of both (FCRA) as dynamic inputs. The FC experiment, forced only with archived forecasting data from ECMWF HRES, serves as a lower benchmark. The RA experiment, exclusively driven by reanalysis (e.g., ERA5) or spatially interpolated observational (e.g., E-OBS) data sources, establishes an upper benchmark for model performance under ideal hindcast conditions (i.e. retrospective simulations using quality-controlled historical data). The FCRA experiment utilized both forecast and reanalysis data for training and testing, representing the optimal data availability scenario.

The domains in the experiments formulate as:

$$\mathcal{D}_{FC} = \{(x_t^{FC}, q_{max,t})\}_{t=1}^T \quad (1)$$

$$\mathcal{D}_{RA} = \{(x_t^{RA}, q_{max,t})\}_{t=1}^T \quad (2)$$

$$\mathcal{D}_{FCRA} = \{(x_t^{FC} \cup x_t^{RA}, q_{max,t})\}_{t=1}^T \quad (3)$$

$\mathcal{D}$... Dataset used in the experiment

x_t ... Meteorological variables at timestep t

$q_{max,t}$... Maximum daily discharge (target variable) at timestep t

2.2.2 Cross-Domain Evaluation

The cross domain evaluation (CD) experiment examined model generalization by training on reanalysis data and testing on ECMWF-HRES forecast data while maintaining 5 identical input variables. The meteorological variables used in this experiment are temperature, dewpoint temperature, precipitation, solar radiation and actual evapotranspiration. This experimental design mirrors the operational framework of classic conceptual hydrological models, where models are typically calibrated using high-quality reanalysis data with subsequently applied real-time forecast inputs during operational usage. By replicating this established modelling paradigm within the LSTM framework, the experiment quantifies the performance shift when transitioning from reanalysis to operationally available forecast data. Our hypothesis here was that if the distributions of the two input data sets $\mathcal{D}_{FC}$ and $\mathcal{D}_{RA}$ are too different, the model performance will decline.



The domains in the experiments formulate as:

$$180 \quad \text{Train on: } \mathcal{D}_{RA} = \{(x_t^{RA}, q_{max,t})\}_{t=1}^T \quad (4)$$

$$\text{Test on: } \mathcal{D}_{FC} = \{(x_t^{FC}, q_{max,t})\}_{t=1}^T \quad (5)$$

2.2.3 Encoder - Decoder LSTM

The Encoder- Decoder LSTM developed by Nearing et al. (2024) consists of two connected LSTMs: one for the hindcast phase forced with historical meteorological reanalysis data (e.g. ERA5) and one for the forecast phase forced with weather forecast data. The two LSTMs are connected by a non-linear handoff network in which the cell state and hidden state from the hindcast are transferred to the forecast LSTM. This architectural design allows the forecast LSTM to learn hydrological states from the hindcast, which could be understood as initial conditions in the model.

Three distinct experiments were implemented using the Encoder- Decoder LSTM architecture to investigate if this dual-LSTM framework can learn and compensate dynamical biases inherent in meteorological forecasts. The first experiment with domain $\mathcal{D}_{ED-LSTM,1}$ implemented the basic encoder-decoder framework where the hindcast LSTM was forced with historical reanalysis data while the forecast LSTM processed meteorological forecast data. The second experiment with domain $\mathcal{D}_{ED-LSTM,2}$ extended this architecture by incorporating past mean daily discharge observations alongside reanalysis data in the hindcast LSTM. The third experiment with domain $\mathcal{D}_{ED-LSTM,3}$ extended $\mathcal{D}_{ED-LSTM,2}$ by additionally forcing the hindcast cell of the model with forecast data. This emulates a setting in which archived forecasting data are used in combination with reanalysis data in the hindcast phase.

The domains in the experiments formulate as:

$$\mathcal{D}_{ED-LSTM,1} = \{(x_{t-s:t-1}^{RA} \cup x_t^{FC}, q_{max,t})\}_{t=s+1}^T \quad (6)$$

$$\mathcal{D}_{ED-LSTM,2} = \{(x_{t-s:t-1}^{RA} \cup q_{max,t-s:t-1} \cup x_t^{FC}, q_{max,t})\}_{t=s+1}^T \quad (7)$$

$$\mathcal{D}_{ED-LSTM,3} = \{(x_{t-s:t-1}^{RA} \cup x_{t-s:t-1}^{FC} \cup q_{max,t-s:t-1} \cup x_t^{FC}, q_{max,t})\}_{t=s+1}^T \quad (8)$$

200 $s \dots$ Sequence length

2.2.4 Sequential Forecast LSTM

The Sequential Forecast LSTM experiment employs a two-phase sequential processing strategy to leverage both reanalysis and operationally available forecast data within a unified framework. The architecture consists of separate embedding networks for hindcast and forecast inputs, a shared LSTM layer and a state transfer mechanism that enables knowledge transfer between processing phases (see Sequential Forecast LSTM in NeuralHydrology, Kratzert et al., 2022). In the first phase, the LSTM processes embedded historical reanalysis data to generate hidden and cell states. The second phase continues LSTM processing with embedded forecast data, initialized with the states from the hindcast phase, ensuring that forecast predictions are informed by contextual information learned from historical patterns. The model generates predictions by concatenating outputs from both phases through a prediction head, with the optimization objective to maximize NSE*. This design enables optimal utilization of reanalysis data for learning hydrological patterns while maintaining operational forecasting capabilities through the principled state transfer mechanism.

Experiment one ($\mathcal{D}_{SEQ LSTM,1}$) used the basic Sequential LSTM framework, with only using reanalysis data in the hindcast phase and forecast data in the forecast phase. The second experiment ($\mathcal{D}_{SEQ LSTM,2}$) added to the first domain mean daily discharge observations alongside reanalysis data in the hindcast phase. In the third experiment ($\mathcal{D}_{SEQ LSTM,3}$), we extended $\mathcal{D}_{SEQ LSTM,2}$ by additionally forcing the hindcast phase of the model with archived forecast data.

The domains in the experiments formulate as:

$$\mathcal{D}_{SEQ LSTM,1} = \{(x_{t-s:t-1}^{RA} \cup x_t^{FC}, q_{max,t})\}_{t=s+1}^T \quad (9)$$



$$\mathcal{D}_{SEQLSTM,2} = \{(x_{t-s:t-1}^{RA} \cup q_{max,t-s:t-1} \cup x_t^{FC}, q_{max,t})\}_{t=s+1}^T \quad (10)$$

$$\mathcal{D}_{SEQLSTM,3} = \{(x_{t-s:t-1}^{RA} \cup x_{t-s:t-1}^{FC} \cup q_{max,t-s:t-1} \cup x_t^{FC}, q_{max,t})\}_{t=s+1}^T \quad (11)$$

2.2.5 Transfer Learning

Transfer Learning (TL) is a machine learning paradigm leveraging gained knowledge from a source domain to improve learning performance in a target domain (Goodfellow et al., 2016). Formally, TL aims to improve the predictive performance on the target domain using knowledge from the source domain, with differences potentially existing in the feature space, data distribution, or learning task between the two domains (Zhuang et al., 2021). TL can be categorized into two primary types based on the relationship between source and target domain: The first is homogeneous transfer learning, where both domains share the same feature space (i.e. using identical meteorological variables and catchment attributes) and have the same marginal probability distributions (Weiss et al., 2016). The second is heterogeneous transfer learning, where the feature spaces differ between domains (Pan and Yang, 2010). For our experiments, we used the heterogeneous transfer learning approach - while the learning task stays the same in the conducted experiments, namely predicting maximum daily discharges at a gauge, the feature spaces and its distributions between forecast (target domain) and reanalysis (source domain) data differs, as evidenced by the violin plots in Appendix B.

In the context of forecast bias reduction, transfer learning is used to leverage knowledge from the less bias-influenced reanalysis source data to improve prediction accuracy when applied to the more bias-prone forecast target data. This approach is particularly relevant in contexts involving hydrometeorological data, where reanalysis data represents a post-processed quality-controlled dataset with reduced systematic errors, while forecast data contains dynamical biases from numerical weather prediction models. By pre-training the temporal encoder (LSTM) on reanalysis data, the model learns hydrological process representations that can subsequently be fine-tuned to accommodate the bias characteristics of forecast inputs, potentially improving the model's ability to correct for systematic forecast errors while maintaining learned temporal dependencies.

The first experiment implemented full weight transfer learning, where all network weights from the baseline $\mathcal{D}_{RA}$ experiment (embedding networks, LSTM and output layers pre-trained on reanalysis data) were used as initialization for a new training phase on forecast data, allowing all parameters to be updated through backpropagation to adapt to the $\mathcal{D}_{FC}$ target domain's characteristics. The second experiment, also based on the weights of $\mathcal{D}_{RA}$ employed selective weight transfer learning, adapting only the embedding network weights while freezing other model parameters, thus preserving learned temporal patterns while allowing adaptation to new input characteristics. The third experiment applied the same selective transfer learning method as the second experiment, but network weights are based on the $\mathcal{D}_{FCRA}$ domain.

The domains in the experiments are given below with source and target domains as well as training objective. All three experiments are based on training either LSTM weights θ , embedding layer weights ϕ , output layer weights ψ or all combined with a NSE* loss function $\mathcal{L}_{NSE}()$.

$$TL_{AllWeights}: \text{source domain } \mathcal{D}_{RA}, \text{ target domain } \mathcal{D}_{FC} \text{ with objective } \arg \min_{\theta, \phi, \psi} \mathcal{L}_{NSE}(\mathcal{D}_{FC}) \quad (12)$$

$$TL_{EmbeddingNet_1}: \text{source domain } \mathcal{D}_{RA}, \text{ target domain } \mathcal{D}_{FC} \text{ with objective } \arg \min_{\phi} \mathcal{L}_{NSE}(\mathcal{D}_{FC}) \quad (13)$$

$$TL_{EmbeddingNet_2}: \text{source domain } \mathcal{D}_{FCRA}, \text{ target domain } \mathcal{D}_{FC} \text{ with objective } \arg \min_{\phi} \mathcal{L}_{NSE}(\mathcal{D}_{FC}) \quad (14)$$

2.2.6 Input Embedding

Input embedding networks serve as pre-processing layers that transform raw meteorological variables into fixed dimensional representations for LSTM processing. The embedding layers enable the model to learn (non-) linear combinations and scaling of input features, potentially capturing complex relationships between meteorological variables that may not be apparent in



their original form (Irani et al., 2025). The embedding transformation is relevant for hydrometeorological applications where variables such as temperature, precipitation and solar radiation may exhibit non-linear interactions that influence runoff generation processes. To investigate the impact of embedding complexity on bias correction performance, we implemented two distinct embedding architectures: a simple embedding consisting of a single fully connected layer with 16 hidden units and tanh activation, and a complex embedding featuring a three-layer network with 30, 20, and 64 hidden units respectively, also using tanh activation functions. The simple embedding provides a lightweight transformation with minimal parameter overhead, while the complex embedding offers greater representational capacity through deeper non-linear transformations.

3 Results and Discussion

All results presented in subsequent sections are visualized as Cumulative Density Functions (CDFs) of Nash-Sutcliffe Efficiency values computed across the 451 study basins, where each point represents the proportion of basins achieving a specified NSE value. This visualization approach enables comprehensive assessment of model performance distribution, revealing not only median performance but also the full range of model behaviour across diverse catchment conditions. The light grey line denotes the forecast-only baseline ($\mathcal{D}_{FC}$) representing the lower performance bound, while the dark grey line denotes the reanalysis-only baseline ($\mathcal{D}_{RA}$) establishing the upper performance bound. These reference curves remain consistent across all figures to facilitate direct comparison between experimental configurations.

3.1 Propagation of Forecast Uncertainty in the Hydrological Model Setting

A fast and straightforward way to analyse the propagation of forecast uncertainty in predicting maximum daily discharge ($q_{\max}$) is the cross- domain evaluation (CD) experiment, depicted in Figure 3. CD reveals degradation in hydrological model performance when transitioning from reanalysis to forecast meteorological forcings, despite five identical input variables. The median Nash-Sutcliffe Efficiency (NSE) decreased from 0.58 to 0.33, representing a 0.25 reduction in model skill. This performance deterioration is accompanied by increased uncertainty, with the NSE standard deviation rising from 0.87 to 1.1, indicating that forecast uncertainty propagates through the hydrological model shifting and broadening the NSE distribution. The mean NSE exhibits an even more pronounced decline (0.44 to 0.19), suggesting increased negative skewness due to extreme poor-performing outliers.

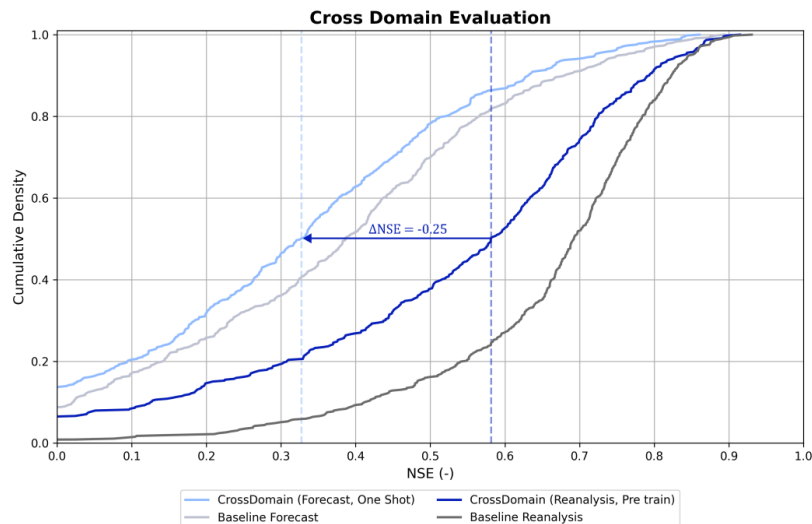


Figure 3. Cumulative Density Function of Nash Sutcliffe Efficiency values for the Cross Domain Evaluation Experiment. The comparison includes the Pre trained model on five meteorological reanalysis variables (dark blue), the One Shot (direct application without fine-tuning) based on the weights of the Pre trained model with equal five meteorological forecasting variables and the baselines (grey). The vertical

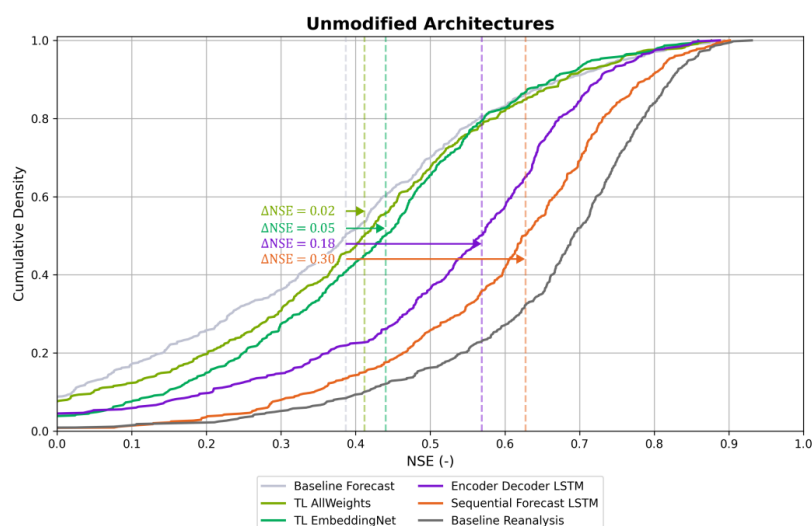


285 dashed lines depict the median NSE for each experiment. The blue arrow shows the performance decrease at median NSE when applying cross domain evaluation.

The underlying cause of this performance decrease can be attributed to the difference in data distributions between reanalysis and forecast datasets (see Appendix B), effectively representing a domain shift problem. Domain shift occurs when the statistical properties of the training data (reanalysis) differ from those of the target data (forecast), violating the fundamental assumption of independent and identically distributed data that underlies machine learning model generalization (Goodfellow et al., 2016; Hosna et al., 2022). Neural networks are particularly susceptible to domain shift as they learn to map input-output relationships based on the specific distributional characteristics of their training data, leading to degraded performance when deployed on data from a different distribution. These results demonstrate that it is not feasible to simply substitute reanalysis data with forecast data in neural network-based hydrological modeling applications, as meteorological forecast uncertainty propagates through the model chain, degrading the representation of catchment processes and creating performance risks for operational hydrological forecasting systems.

3.2 Performance Analysis Across Unmodified Architectures

To address the domain shift challenges identified in Section 3.1, we evaluated different neural network architectures and training techniques, presenting here the optimal configurations from each experimental setup.



300 **Figure 4.** Cumulative Density Function of Nash Sutcliffe Efficiency values for the Unmodified Architectures Experiment. The experiments include the baselines (grey), comprehensive transfer learning with a finetuning of the embedding and LSTM weights (TL AllWeights), selective transfer learning with only finetuning the embedding weights while the LSTM was freezed (TL EmbeddingNet), Encoder-Decoder LSTM (purple) and the Sequential Forecast LSTM (orange). The vertical dashed lines depict the median NSE for each experiment. The arrows indicate the median Δ NSE values relative to the baseline forecast experiment.

310 Transfer learning was implemented through two contrasting approaches: comprehensive parameter updating, where the entire network was retrained on forecast data (TL AllWeights), and selective embedding retraining, where only the input embedding layers were fine-tuned while maintaining the pre-trained weights of the deeper network components (TL EmbeddingNet). Both experiments utilized the reanalysis baseline with a complex embedding layer as the starting point. The selective embedding approach (TL EmbeddingNet) achieved higher performance, with a median NSE of 0.44, representing a 0.11 improvement over the cross-domain baseline (0.33). The advantage of selective retraining becomes evident when comparing the two transfer learning strategies: TL EmbeddingNet shows an improvement in the 10th percentile (0.15) compared to TL AllWeights (0.05), indicating that the retraining approach is especially effective for poorly performing basins. The enhanced performance stems from the model's ability to leverage robust hydrometeorological relationships learned from reanalysis data while adapting the



input representation to forecast data characteristics. Reanalysis products provide more complete and physically consistent atmospheric descriptions through data assimilation, enabling the model to learn generalized process representations that are subsequently refined during fine-tuning on forecast data.

The Encoder-Decoder LSTM architecture demonstrated performance improvements over the transfer learning approaches, achieving a median NSE of 0.57. This architecture showed particular strength in the upper performance range, with the 75th and 90th percentiles reaching 0.65 and 0.73, respectively.

The Sequential Forecast LSTM achieved the highest overall performance among the unmodified forecast-based configurations, with a median NSE of 0.63 and notably consistent results across all percentiles. The model demonstrated higher stability compared to other approaches, evidenced by the low standard deviation (0.52). Among the unmodified architectures, the Sequential LSTM is best able to efficiently reduce forecasting bias. We attribute this to the sequential data processing architecture of this special LSTM type, which processes meteorological forecast inputs in temporal order and allows for the gradual correction of forecast biases through the propagation of both hidden states and cell states. The cell state serves as long-term memory that can selectively retain or forget information across time steps, while the hidden state captures current relevant information, enabling the model to maintain both short-term adaptations to recent forecast patterns and long-term memory of systematic biases. Unlike the Encoder- Decoder architecture transforming hindcast states through a fixed handoff network before initiating forecast processing, the Sequential LSTM maintains continuous state evolution from hindcast to forecast, preserving temporal patterns without disruption and potentially enabling better compensation for systematic errors that emerge at the forecast transition.

3.3 The role of integrating archived forecasts in the hindcast phase

Given the performance degradation observed when transitioning from reanalysis to forecast data, we investigated whether training models with a combination of reanalysis and archived forecast data could improve forecast performance. In short, the integration of archived forecasts in the hindcast phase demonstrates limited effectiveness across all tested architectures.

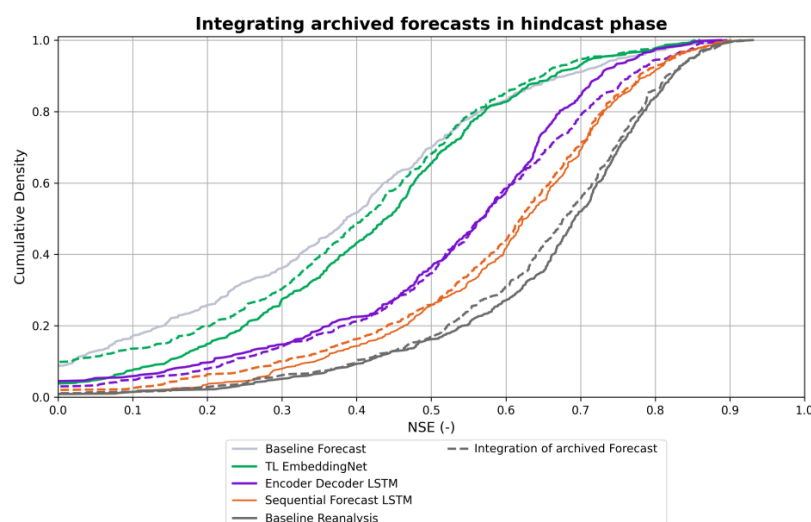


Figure 5. Cumulative Density Function of Nash Sutcliffe Efficiency values for experiments integrating archived forecasts in the hindcast phase. The experiments include the baselines (grey), selective transfer learning with only finetuning the embedding weights while the LSTM was frozen (TL EmbeddingNet), Encoder- Decoder LSTM (purple) and the Sequential Forecast LSTM (orange). Solid lines indicate the experiments with solely reanalysis data in the hindcast phase, while dashed lines display the combination of archived forecasts and reanalysis in the hindcast phase.

In Figure 4 it is evident that in transfer learning only fine-tuning the embedding net led to higher forecast skill, leading us to only focus in this experimental setup on the selective transfer learning method. As previously discussed, the TL EmbeddingNet



approach pretrained exclusively on reanalysis data achieved a median NSE of 0.44. In contrast, when the same transfer learning architecture is pretrained on combined forecast and reanalysis data, the results show deteriorated performance compared to the reanalysis-only pretraining, with a median NSE dropping to 0.41. This configuration exhibits increased variability (standard deviation of 2.29) while the 10th percentile performance remains poor at 0.02, failing to achieve the improvement (0.15) observed with reanalysis-only pretraining. The mean performance also deteriorates significantly (0.21 vs 0.35 for reanalysis-only), indicating the introduction of more negative outliers.

For the Encoder-Decoder LSTM, incorporating archived forecasts yields marginal improvements in the upper performance percentiles, with the 75th and 90th percentiles increasing from 0.65 to 0.69 and 0.73 to 0.77, respectively, while the 10th percentile improves slightly from 0.21 to 0.24. However, these modest gains are accompanied by increased variability (standard deviation increases from 8.03 to 9.33) while the median NSE remains unchanged at 0.57.

The Sequential Forecast LSTM shows no meaningful benefit from archived forecast integration, with the median NSE decreasing marginally from 0.63 to 0.62. More critically, this architecture experiences an increase in variability (standard deviation from 0.52 to 6.28) and the mean performance drops from 0.57 to 0.19, indicating the introduction of numerous extreme negative outliers that significantly compromise model reliability.

The results demonstrate that integrating archived forecasts during the hindcast phase does not provide meaningful performance improvements for either architecture. The approach either yields negligible benefits while increasing instability (Encoder-Decoder LSTM) or actively degrades performance (Transfer learning, Sequential LSTM), indicating that this strategy is not effective for addressing domain shift challenges in hydrological forecasting applications.

3.4 The role of integrating past discharge in the training domain

Previous studies have consistently focused on modeling ungauged basins (Kratzert et al., 2019a; Nearing et al., 2024). However, we argue that when near real-time discharge data are available, as is the case for most gauging stations across the LamaH catchments in Central Europe, the incorporation of discharge observations positively influences forecast accuracy. For example, in the Austrian LamaH basins, discharge data from the eHYD platform (<https://ehyd.gv.at>) are available with a time delay of only two hours, making the integration of recent discharge observations operationally feasible for real-time forecasting applications.

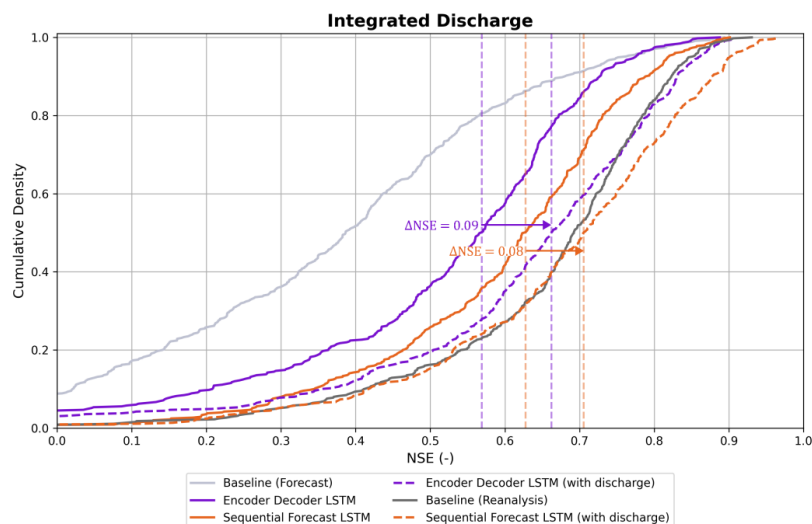


Figure 6. Cumulative Density Function of Nash Sutcliffe Efficiency values for experiments integrating past discharge in the hindcast phase. The experiments include the baselines (grey), Encoder-Decoder LSTM (purple) and Sequential Forecast LSTM (orange). Solid lines depict the experiments without discharge, while the dashed lines show experiments with discharge in the hindcast phase. The arrows represent the prediction accuracy increase as median ΔNSE when integrating discharge in the hindcast phase of the models.



375 The integration of past discharge observations demonstrates performance improvements across both tested architectures. For
the Encoder-Decoder LSTM, incorporating discharge data yields significant gains, with the median NSE increasing from 0.57
to 0.66. The improvement is particularly pronounced in the lower percentiles, with the 10th percentile rising from 0.21 to 0.37,
indicating considerably better performance for poorly performing basins. The 75th and 90th percentiles also show notable
improvements (0.65 to 0.77 and 0.73 to 0.83, respectively), while the variability decreases slightly (standard deviation from
380 8.03 to 6.68).

The Sequential Forecast LSTM exhibits even more impressive improvements when discharge is integrated, achieving a median
NSE of 0.71 compared to 0.63 without discharge. This represents the highest performance among all forecast-based
configurations and even exceeds the reanalysis baseline performance (0.69). The 10th percentile shows substantial
improvement (0.35 to 0.42), while the upper percentiles reach 0.81 and 0.88 for the 75th and 90th percentiles, also exceeding
385 the reanalysis baseline performance in these ranges.

Additional transfer learning experiments were conducted to investigate whether discharge information learned during pre-
training could be effectively transferred to discharge-free operational scenarios. These experiments included domain adaption
configurations where discharge was incorporated during the pre-training phase but excluded during fine-tuning, as well as
setups using the Sequential Forecast LSTM with discharge in the source domain and without discharge in the target domain.
390 All transfer learning approaches consistently resulted in performance deterioration and failed to yield improvements that would
justify the computational overhead of the transfer learning process. While these experiments demonstrated that information
extraction from LSTM cell states is feasible, the learned discharge-related representations could not adequately compensate
for the absence of direct discharge observations during operational forecasting. The experiments with discharge incorporated
directly into the training data consistently outperformed all transfer learning alternatives, indicating that real-time discharge
395 integration provides irreplaceable benefits that cannot be effectively substituted through knowledge transfer mechanisms.

3.5 The role of embedding complexity

During experimenting we experienced that the embedding complexity plays an important role in reducing the forecast induced
bias in runoff predictions. This circumstance has led us to create this experimental setup, in which a simple (linear) and a
complex (non-linear) embedding were generated for all architectures used in the previous experimental settings. In brief, a
400 similar pattern can be observed for all architectures except transfer learning: the more complex the input embedding, the lower
the prediction performance.

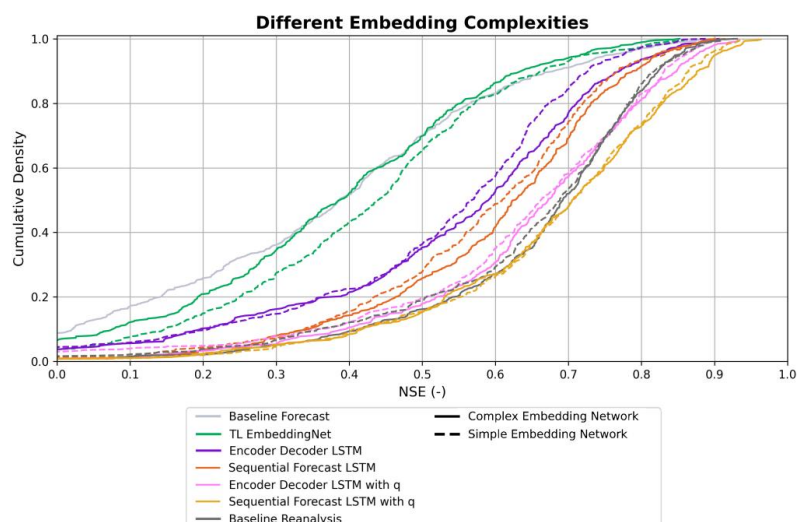


Figure 7. Cumulative Density Function of Nash Sutcliffe Efficiency values for experiments with different embedding complexities. The experiments include the baselines (grey), selective transfer learning with only finetuning the embedding weights while the LSTM was frozen (TL EmbeddingNet), Encoder- Decoder LSTM without discharge (purple), Encoder- Decoder LSTM with discharge (pink), Sequential Forecast LSTM without discharge (orange) and Sequential Forecast LSTM with discharge (yellow). Solid lines depict the experiments without simple linear embedding, while the dashed lines show experiments with complex non- linear embedding networks.

While the transfer learning experiments did not achieve the high NSE values of the best-performing forecasting approach (Sequential LSTM), notable relationships emerged between embedding complexity and prediction accuracy. In the baseline, encoder-decoder and sequential LSTM experiments, simple linear embedding networks produced slightly higher simulation performance compared to more complex embedding architectures. This pattern reversed in transfer learning experiments, where increased embedding complexity led to improved prediction results. The transfer learning procedure consisted of transferring weights from an LSTM model pre-trained on reanalysis data to a new network architecture with modified input embeddings, where only the embedding weights underwent retraining while the remaining LSTM parameters remained frozen.

The minimal better performance of simple embeddings in non-transfer learning experiments can be attributed to the tendency of complex embedding networks to overfit when trained on the available data. The number of trainable parameters in complex embeddings exceeds the optimal ratio relative to available training data, causing the model to learn specific noise patterns in the training data rather than generalizable hydrological patterns. This means simpler embeddings capture the relevant input patterns more effectively without introducing unnecessary model complexity comprising generalization.

The improved results with increased embedding complexity in transfer learning can be explained by the requirement for flexible, non-linear transformations necessary for effective domain adaptation. Since only the embedding weights are trained while the all other network parameters remain frozen, the embedding layer must perform all adaptation work between the target domain and the representations pre-trained on reanalysis data. More complex architectures can perform richer and more domain-specific feature extraction, compensating for the discrepancy between source and target domains through more expressive input transformations. The contrasting performance patterns between transfer learning and standard training suggest that the optimal embedding complexity depends on whether the model parameters are trained from scratch or adapted from pre-trained weights, though the precise mechanisms underlying this relationship warrant further investigation.

3.6 Limitations and Future Directions

This study is subject to several limitations. The experiments were conducted on the Extended LamaH-CE dataset, which is restricted to Central Europe; the transferability of the results to other hydroclimatic regions remains to be tested. While



discharge observations were shown to substantially improve performance, they are not error-free, and their availability cannot be guaranteed in ungauged or poorly monitored catchments.

In addition, we only evaluated existing LSTM-based architectures and a modeling setup with spatially lumped data and daily maximum discharges. Applying these approaches in fully distributed settings with higher temporal resolution forecasts may
 435 yield different, potentially more pronounced results. The integration of reanalysis data in the hindcast phase, while useful for training, is also constrained by data latency, which is often critical for operational use, although future improvements in near-real-time reanalysis may alleviate this.

Addressing these challenges—testing transferability, reducing dependence on delayed or unavailable inputs, and extending both the spatial and temporal resolution of experiments—will be essential to further investigate and advance robust,
 440 operationally viable solutions for flood forecasting.

4 Conclusion and Outlook

The operational deployment of deep learning based flood forecasting models faces fundamental challenges when transitioning from high-quality reanalysis to meteorological forecast data, with domain shift between these data sources leading to model performance degradation. While previous approaches have focused on bias-correcting meteorological inputs through statistical
 445 methods (Lenderink et al., 2007) or machine learning techniques applied to precipitation forecasts (Ko et al., 2020; Zhang et al., 2020), these methods rely on comparing forecasts with meteorological observations that themselves contain uncertainties (Bárdossy et al., 2022). This study addressed the challenge through systematic evaluation of Long Short-Term Memory architectures and training techniques that learn bias correction directly from the more reliable discharge observations, following the paradigm suggested by Kirchner (2009) of using river discharge as the primary constraint. Our experiments
 450 across 451 Central European catchments demonstrated that appropriate neural network designs can transform the domain shift problem from a major obstacle into a learnable pattern correction task. Sequential Forecast LSTM architectures, when combining meteorological hindcast data with past discharge observations, provided the most effective framework for mitigating forecast-induced biases. This configuration achieved a median NSE of 0.71, surpassing even the reanalysis baseline simulation and establishing discharge integration, if data are available in near real time as in the LamaH domain, as a critical
 455 component for operational forecast accuracy.

To quantify the bias propagation caused by the domain shift, we conducted cross-domain evaluation revealing performance deterioration when reanalysis-trained models were applied to forecast inputs. In this setting, the median Nash-Sutcliffe Efficiency decreased from 0.58 to 0.33, representing a 0.25 reduction in model skill. This performance degradation stems from fundamental differences in data distributions between reanalysis and forecast datasets, violating the assumption of identically
 460 distributed training and testing data that underlies machine learning model generalization (Goodfellow et al., 2016). Among the tested neural network architectures, the Sequential Forecast LSTM demonstrated superior performance for operational forecasting applications. This architecture achieved a median NSE of 0.63 with notable stability (standard deviation of 0.52) and maintained reasonable performance across all percentiles. The sequential processing approach enables gradual correction of forecast biases through continuous state evolution from hindcast to forecast phases, preserving temporal patterns without
 465 the disruption introduced by fixed handoff networks in Encoder-Decoder architectures. Transfer learning approaches, despite theoretical advantages for domain adaptation, achieved only modest improvements (NSE 0.44), while the incorporation of archived forecasts in training failed to provide consistent benefits and often increased model instability. The relationship between embedding complexity and performance varied systematically: simpler embeddings performed better in standard training contexts, while complex embeddings showed advantages only in transfer learning scenarios where flexible input
 470 transformations were required for domain adaptation.



These findings carry important implications for operational flood forecasting system design. The high performance of Sequential Forecast LSTM architectures indicates that operational systems should prioritize continuous state transfer mechanisms maintaining temporal dependencies across the hindcast-forecast phases rather than treating these phases as disconnected processes. The substantial improvements from discharge integration align with operational capabilities in many monitored systems, such as in hydro power plants or the Austrian eHYD platform where observations are available with two-hour latency, making real-time integration feasible. The ability to learn bias correction patterns directly from the combined meteorological-hydrological data space eliminates the need for separate pre-processing steps, whether meteorological bias correction (Hess, 2020; Han et al., 2021) or streamflow post-processing methods like the GloFAS-specific approach of Hunt et al. (2022), reducing computational overhead and possibly potential error propagation.

Future developments should focus on adaptive architectures that can dynamically leverage discharge observations when available while maintaining robust performance in ungauged settings through reanalysis-only hindcast processing. Such unified frameworks would enable seamless deployment across both gauged and ungauged basins within the same operational system, automatically adjusting to data availability in real-time. However, sensor failures should also be taken into account here, and the training methods proposed by Gauch et al. (2025) should be applied. While our experiments were limited to ECMWF HRES archived forecasts due to data availability, combining multiple forecasts from different sources could also be promising as it could capture forecast uncertainty more comprehensively and enable the model to learn source-specific bias patterns, as with the integration of multiple meteorological data in a simulation setting (Kratzert et al., 2021). The Sequential Forecast LSTM's bias correction capabilities at 24-hour lead times provide a strong foundation for multi-day forecasting applications, where learning lead-time dependent bias patterns could improve medium-range flood predictions that are crucial for early warning systems and emergency preparedness. The demonstrated ability of LSTM architectures and training techniques to transform the domain shift challenge into a learnable bias correction problem, combined with increasing availability of real-time hydrological observations, establishes a pathway toward operational flood forecasting systems that can maintain predictive skill despite the inherent uncertainties coming from numerical weather predictions.

Code and data availability. All experiments have been conducted with a forked version of the NeuralHydrology library (Kratzert et al., 2022), available at github.com/conestone/neuralhydrology. The Extended LamaH-CE dataset is available at zenodo.org/records/17119635 (Konold et al., 2025b). The code to create the analysis and figures is available at github.com/conestone/biascast. All trained models with its configuration files and saved weights are available at 10.5281/zenodo.17241922 (Konold et al., 2025a).

500



Appendix A. Dynamic and Static Forcings of the Models

Variable	Description	Unit	Source Product	Source
ERA5L_2m_temp_max	2m above earth surface max air temperature	°C	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_2m_temp_mean	2m above earth surface mean air temperature	°C	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_2m_temp_min	2m above earth surface min air temperature	°C	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_2m_dp_temp_max	2m above earth surface max dewpoint temperature	°C	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_2m_dp_temp_mean	2m above earth surface mean dewpoint temperature	°C	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_2m_dp_temp_min	2m above earth surface min dewpoint temperature	°C	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_10m_wind_u	Eastwards wind speed 10m above earth surface	m/s	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_10m_wind_v	Northwards wind speed 10m above earth surface	m/s	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_fcst_alb	Forecast albedo	-	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_lai_high_veg	Leaf Area Index for high vegetation type	m ² /m ²	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_lai_low_veg	Leaf Area Index for low vegetation type	m ² /m ²	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_swe	Snow Water Equivalent	mm	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_surf_net_solar_rad_max	Max amount of solar radiation reaching the Earth's surface minus the amount reflected by the Earth's surface	W/m ²	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_surf_net_solar_rad_mean	Mean amount of solar radiation reaching the Earth's surface minus the amount reflected by the Earth's surface	W/m ²	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_surf_net_therm_rad_max	Maximum net thermal radiation at the Earth's surface;	W/m ²	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_surf_net_therm_rad_mean	Mean net thermal radiation at the Earth's surface;	W/m ²	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_surf_press	Surface pressure	Pa	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_total_et	Total evapotranspiration	mm	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_prec	Total precipitation	mm	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_volsw_123	Fraction of water from 0 to 100 cm depth (topsoil)	m ³ /m ³	ERA5Land	Muñoz-Sabater et al., 2021
ERA5L_volsw_4	Fraction of water from 100 to 289 cm depth (subsoil)	m ³ /m ³	ERA5Land	Muñoz-Sabater et al., 2021
EOBS_tg	2m above earth surface mean daily air temperature	°C	E-OBS	Cornes et al., 2018
EOBS_tn	2m above earth surface min daily air temperature	°C	E-OBS	Cornes et al., 2018
EOBS_tx	2m above earth surface max daily air temperature	°C	E-OBS	Cornes et al., 2018
EOBS_rr	Total precipitation	mm	E-OBS	Cornes et al., 2018
EOBS_pp	Mean sea level pressure	hPa	E-OBS	Cornes et al., 2018
EOBS_fg	Mean wind speed at 10m height	m/s	E-OBS	Cornes et al., 2018
EOBS_qq	Solar radiation at earth's surface	W/m ²	E-OBS	Cornes et al., 2018
MSWEP_RR	Total precipitation	mm	MSWEP	Beck et al., 2019
GLEAM_ETA	Actual evapotranspiration	mm	GLEAM	Miralles et al., 2011
GLEAM_ETP	Potential evapotranspiration	mm	GLEAM	Miralles et al., 2011
ECMWF_t2m	Forecasted 2m above earth surface mean air temperature	°C	ECMWF-HRES	ECMWF, 2025
ECMWF_d2m	Forecasted 2m above earth surface mean dewpoint temperature	°C	ECMWF-HRES	ECMWF, 2025
ECMWF_ssrd	Forecasted solar radiation at earth's surface	J/m ²	ECMWF-HRES	ECMWF, 2025
ECMWF_tp	Forecasted total precipitation	mm	ECMWF-HRES	ECMWF, 2025
ECMWF_e	Forecasted total actual evapotranspiration	mm	ECMWF-HRES	ECMWF, 2025

Table A1. Meteorological Variables in the Extended LamaH-CE data set used for the conducted experiments



Attribute	Description	Unit
area_calc	Calculated basin area	km ²
elev_mean	Mean catchment elevation	m a.s.l
elev_med	Median catchment elevation	m a.s.l
elev_std	Standard deviation of elevation in catchment	m a.s.l
elev_ran	Range of catchment elevation (max – min elev.)	m a.s.l
slope_mean	Mean catchment slope	m/km
mvert_dist	Horizontal distance from the farthest point of the catchment to the corresponding gauge (length axis)	km ²
mvert_ang	Angle between the north direction and connection from farthest point of catchment to the corresponding gauge (length axis)	degree
elon_ratio	Elongation ratio between the diameter D of an equivalent circle and the area of the catchment area to its length L	-
strm_dens	Stream density	km/km ²
p_mean	Mean daily precipitation	mm/day
et0_mean	Mean daily reference evapotranspiration	mm/day
eta_mean	Mean daily total evapotranspiration	mm/day
arid_1	Aridity, computed as the ratio of mean et0_mean and p_mean	-
arid_2	Reciprocal value of aridity index	-
p_season	Seasonality and timing of precipitation (estimated using sine curves) to represent the annual precipitation cycles	-
frac_snow	Fraction of precipitation falling as snow	-
hi_prec_fr	Frequency of high-precipitation days	day/year
hi_prec_du	Mean duration of high-precipitation events	day
lo_prec_fr	Frequency of dry days	day/year
lo_prec_du	Mean duration of dry periods	day
lc_dom	Three-digit short code of dominant land cover class	-
agr_fra	Fraction of agricultural areas	-
bare_fra	Fraction of bare areas	-
forest_fra	Fraction of forest areas	-
lake_fra	Fraction of natural or artificial water bodies with all-season water filling	-
urban_fra	Fraction of areas mainly occupied by buildings including their connected areas	-
lai_max	Maximum monthly mean of one-sided leaf area index	m ² /m ²
lai_diff	Difference between maximum and minimum monthly mean of one-sided leaf area index	m ² /m ²
ndvi_max	Maximum monthly mean of NDVI	-
ndvi_min	Minimum monthly mean of NDVI	-
gvf_max	Maximum monthly mean of the green vegetation fraction	-
gvf_diff	Difference between the maximum and minimum monthly mean of the green vegetation fraction	-

505 **Table A2.** Static catchment attributes from the Extended LamaH-CE data set used for the conducted experiments



Appendix B. Variability across input data sets

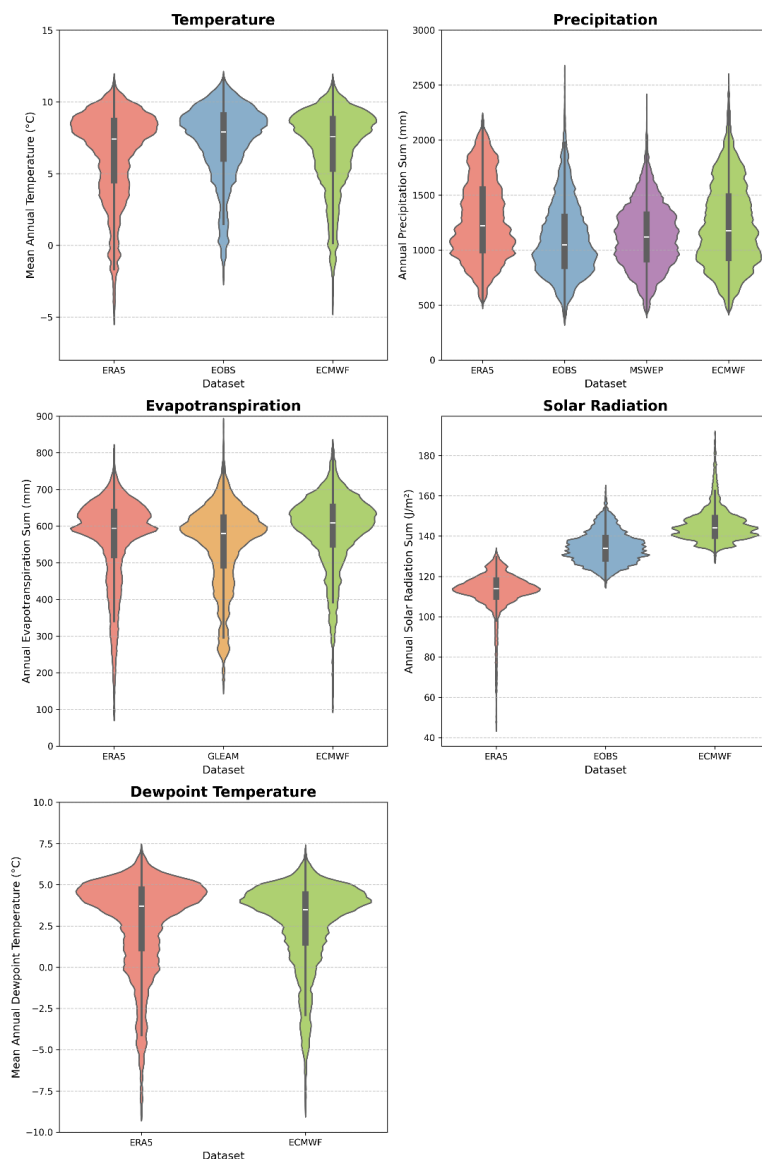


Figure B1. Variability across input data sets displayed as violin plots for annually aggregated meteorological variables Temperature, Precipitation, Actual Evapotranspiration, Solar Radiation and Dewpoint Temperature. The violin colours belong always to a certain data product: red- ERA5Land, blue: E-OBS, green: ECMWF-HRES, purple: MSWEP. The dark grey box in the violin shows a boxplot with the bold part depicting the interquartile range and the white line indicating the median.



Appendix C. Evaluation Metrics

515 The model performance was evaluated by using the non basin specific NSE* of Kratzert et al. (2019b) which is based on the Nash and Sutcliffe Efficiency (NSE; Nash and Sutcliffe, 1970). In comparison to former studies (Kratzert et al., 2019a, b, 2021; Nearing et al., 2024), we conducted no cross validation across spatial units, since we did not focus on ungauged basins. Instead, we employed a temporal split validation approach, where the available time series data for each catchment was divided into training, validation, and testing periods to ensure robust model evaluation.

$$520 \quad NSE = 1 - \frac{\sum_{i=1}^n (y_{obs,i} - y_{sim,i})^2}{\sum_{i=1}^n (y_{obs,i} - \bar{y}_{obs})^2} \quad (C1)$$

$$NSE^* = \frac{1}{B} \sum_{b=1}^B \sum_{n=1}^N \frac{(\bar{y}_n - y_n)^2}{(s(b) + \epsilon)^2} \quad (C2)$$

Appendix D. Hyperparameter Optimization

Hyperparameter tuning is a critical component in the bias correction of meteorological forecasting data using Long Short-Term Memory (LSTM) networks, as it directly influences the model's ability to learn complex temporal patterns and correct systematic biases in forecast inputs. Given the nonlinear and dynamic nature of meteorological variables, appropriate selection of hyperparameters such as learning rate, sequence length, number of hidden units, dropout rate, and batch size, is essential to ensure that the LSTM model generalizes well without overfitting to noise or underfitting relevant signals. In the context of bias correction, the model must not only capture historical dependencies in the forecast errors but also effectively differentiate between genuine atmospheric variability and persistent model biases. Without careful tuning, the LSTM may fail to correct biases accurately, particularly under extreme events or seasonal transitions.

To this end, we used Bayesian optimization as described by Peter I. Frazier (2018). This search algorithm fits a Gaussian process to the observed hyperparameter-performance pairs to estimate performance on yet-untested parameter settings. We use the expected improvement as acquisition function, which is used for selecting the next set of hyperparameters to test. This approach efficiently identifies good hyperparameters, especially in large search spaces.

535 For our models, we spanned the search space over the hidden layer sizes (64, 128, 256 units), output dropout rates (0.1, 0.2, 0.3), variance of Gaussian noise applied to the discharge values (0.001, 0.01, 0.1), and batch sizes (64, 128, 256) and limited the number of iterations to 100. To obtain the result for each setting, we trained a model for 30 epochs, an initial learning rate of 0.001 and a cosine annealing schedule ($T_{max}=30$, $\eta_{min}=1e-5$), stopping early if the model did not improve the evaluation metric by more than 0.005 in the last 5 epochs. To estimate the performance of the model we used the basin averaged Nash-Sutcliffe efficiency (NSE*) metric and evaluated it on the validation period.

NSE* with hyperparameter optimization:

$$\lambda = \arg \max_{\lambda} NSE_{median}(\mathcal{D}_{val}; \lambda) \quad (D1)$$

where

$$\lambda = \{d_h, p_{dropout}, \sigma_{noise}, b_{size}\} \quad (D2)$$

545 represents the optimal output.



Appendix E. Long-Short Term Memory Network

Long term information is stored in the cell state (c_t), short term information in the hidden state (h_t) and the information flow is controlled by the so- called gating mechanisms (Hochreiter and Schmidhuber, 1997). The input gate (i_t) determines how much of the current input and previous hidden state contributes to updating the cell state. The forget gate (f_t) regulates which parts of the previous cell state should be retained or discarded, allowing the model to reset its memory when necessary. The output gate (o_t) defines how much of the updated cell state is exposed to the next time step via the hidden state (Gers et al., 1999). This architecture allows LSTMs to retain relevant information over longer time periods and capture temporal dependencies in input data.

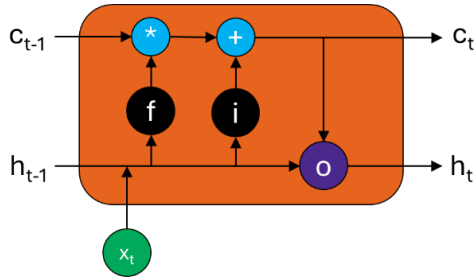


Figure E1. LSTM cell

$$f_t = \sigma(W_f x_t + U_f h_{t-1} + b_f) \quad (E1)$$

$$\tilde{c}_t = \tanh(W_c x_t + U_c h_{t-1} + b_c) \quad (E2)$$

$$i_t = \sigma(W_i x_t + U_i h_{t-1} + b_i) \quad (E3)$$

$$c_t = f_t \odot c_{t-1} + i_t \odot \tilde{c}_t \quad (E4)$$

$$o_t = \sigma(W_o x_t + U_o h_{t-1} + b_o) \quad (E5)$$

$$h_t = \tanh(c_t) \odot o_t \quad (E6)$$

Appendix F. Computational Resources

All conducted experiments were trained on a NVIDIA RTX4090 graphics processing unit, with wall times varying between several minutes to approximately one hour for one model run, depending on the size of the input vector in the model and the model architecture itself. Although it is common practice in hyperparameter optimisation to run the same settings three times with different seedings, we have only run the tuning with one seed at a time due to computational constraints.



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