

# An improved Bayesian inversion to estimate daily NO<sub>x</sub> emissions of Paris from TROPOMI NO<sub>2</sub> observations between 2018-2023

Alba Mols<sup>1</sup>, Klaas Folkert Boersma<sup>1,2</sup>, Hugo Denier van der Gon<sup>3</sup>, and Maarten Krol<sup>1,4</sup>

<sup>1</sup>Wageningen University, Meteorological and Air Quality department, Wageningen, the Netherlands

<sup>2</sup>Royal Netherlands Meteorological Institute (KNMI), De Bilt, the Netherlands

<sup>3</sup>Department of Climate, Air and Sustainability, TNO, Utrecht, the Netherlands

<sup>4</sup>Institute for Marine and Atmospheric Research Utrecht (IMAU), Utrecht University, Utrecht, the Netherlands

**Correspondence:** Folkert Boersma (folkert.boersma@wur.nl)

**Abstract.** We present a comprehensive quantification of daily NO<sub>x</sub> emissions from Paris using an inverse analysis of tropospheric NO<sub>2</sub> columns measured by the Tropospheric Monitoring Instrument (TROPOMI) over a 5-year period (May 2018 - August 2023). Our analysis leverages a superposition column model that captures the relationship between the increase in NO<sub>2</sub> with distance over an urban source region to underlying NO<sub>x</sub> emissions, accounting for chemical transformations and wind in the urban boundary layer. To evaluate the robustness of the superposition column model, we tested it against high-resolution (300 m) Large Eddy Simulations (LES) using MicroHH, [a computational fluid dynamics model](#), with atmospheric chemistry, confirming that the model's simplifying assumptions introduce uncertainties below 10%. Building on this foundation, we develop a new Bayesian inversion method that incorporates prior knowledge on NO<sub>x</sub> emissions and lifetimes and accounts for model and prior uncertainties. Compared to a previous look-up table approach, which relied on least-squares minimization without prior constraints, the Bayesian method demonstrated superior performance. In controlled tests, it reproduced known NO<sub>x</sub> emissions within 5%. Applying Bayesian inversion to TROPOMI data in Paris, we observed a significant reduction in NO<sub>x</sub> emissions from 44 mol s<sup>-1</sup> in 2018 to 32 mol s<sup>-1</sup> in 2023, representing a ~~4827~~ % decrease. This decline exceeds the 12% reduction predicted by the TNO-MACC-III bottom-up inventory, indicating limited accuracy of current inventories. Seasonal analysis revealed higher posterior emissions in winter, possibly highlighting the role of residential heating or vehicle cold starts, which may be underrepresented in bottom-up estimates. Our improved Bayesian framework delivers accurate NO<sub>x</sub> emission estimates that align well with independent data sets. This approach provides a valuable tool for monitoring urban NO<sub>x</sub> emissions and assessing the efficacy of air quality policies.

## 1 Introduction

Nitrogen oxides (NO<sub>x</sub>=NO+NO<sub>2</sub>) are major air pollutants which are central to the chemistry of the troposphere, and which have negative impacts on human health and the environment (e.g. Boningari and Smirniotis (2016)). In urban regions NO<sub>x</sub> is mainly emitted to the atmosphere as a result of the burning of fossil fuels, particularly in combustion engines. In the EU, the largest contributor to NO<sub>x</sub> emissions is the transport sector (40%), followed by energy production and distribution (16%), and the commercial, institutional and households sectors (15%) (EEA, 2019). At daytime, nitrogen oxides are short-lived, on the

order of 1-12 hours (Stavrakou et al., 2013), because  $\text{NO}_2$  is quickly oxidized by reaction with the hydroxyl radical ( $\text{OH}$ ) to 25 nitric acid ( $\text{HNO}_3$ ), which is deposited by precipitation. Due to its high water solubility,  $\text{HNO}_3$  is efficiently removed from the atmosphere, primarily through precipitation and direct deposition onto surfaces (Seinfeld and Pandis, 2016).

Besides being a toxic gas itself,  $\text{NO}_2$  also has secondary effects via its contribution to photochemical ozone production (Seinfeld and Pandis, 2016; Jacob, 1999; Visser et al., 2019), its influence on the formation of aerosols (Yan et al., 2020), and its contribution to eutrophication via the deposition of  $\text{HNO}_3$  to ecosystems (e.g. Vitousek et al. (1997); Erisman and Draaijers 30 (1995)).

To reduce the negative effects of  $\text{NO}_x$ , the EU maintains a limit value for average annual surface air  $\text{NO}_2$  concentrations of  $40 \mu\text{g m}^{-3}$ . Currently, air quality in most European cities complies (EEA, 2022). Nevertheless,  $\text{NO}_x$  pollution remains a significant health concern for Europeans, especially in urban areas, as daily guidelines set by the World Health Organisation (WHO) of  $25 \mu\text{g m}^{-3}$  (WHO, 2021) continue to be frequently exceeded (EEA, 2022).

35 Monitoring of  $\text{NO}_x$  emissions typically relies on bottom-up inventories, which are uncertain due to their reliance on emission factors, extrapolations and activity assumptions. Uncertainties in bottom-up emissions vary with location, and are estimated to be typically more than 30% (Kuenen et al., 2014). Satellite measurements of  $\text{NO}_2$  offer a useful tool for top-down inverse modelling of  $\text{NO}_x$  emissions, providing more insights into  $\text{NO}_x$  sources and distributions. However, inverse modelling is also subject to assumptions, such as uncertainties of  $\text{NO}_x$  lifetimes in the lower atmosphere (Stavrakou et al., 2013), which can 40 significantly influence the accuracy of top-down emission flux estimates.

Research and refinement of inversion methods for estimating  $\text{NO}_x$  emissions and lifetimes are crucial, especially for initiatives like the Copernicus CO2M mission (Sierk et al., 2021), which will utilize  $\text{NO}_2$  plumes to enhance  $\text{CO}_2$  monitoring by pinpointing emissions more accurately. Several studies have quantified  $\text{NO}_x$  emissions based on satellite  $\text{NO}_2$  retrievals, by analyzing downwind plumes of  $\text{NO}_2$  from large sources, using inverse modeling computations with atmospheric chemical 45 transport models (e.g. Brioude et al. (2013); Cheng et al. (2021); Kurokawa et al. (2009); Krol et al. (2024)). CTMs (e.g. Brioude et al. (2013); Cheng et al. (2021); Kurokawa et al. (2009); Krol et al. (2024); Zhu et al. (2022)). However, because CTMs can have accessibility issues and require substantial computational resources, alternative approaches that do not depend on CTMs have been developed and utilized to estimate  $\text{NO}_x$  emissions and lifetimes. (e.g. de Foy et al. (2014), Beirle et al. (2011) ).

50 Beirle et al. (2011) first presented a method to infer  $\text{NO}_x$  emissions from strong isolated sources, by averaging satellite  $\text{NO}_2$  plumes with similar wind direction. Building upon this concept, Lorente et al. (2019) presented a simple superposition column model that uses  $\text{NO}_2$  retrievals over Paris of the TROPOspheric Monitoring Instrument (TROPOMI), combined with domain-average information about the wind speed, wind direction and  $\text{OH}$  concentrations in the boundary layer, to estimate urban  $\text{NO}_x$  emissions and lifetimes without the need for complex inverse modelling computations. This approach allows for 55 day-to-day emission estimates under cloud-free conditions, offering the potential for continuous emission estimations over a long period of time. Zhang et al. (2022) expanded this model framework to estimate the  $\text{NO}_x$  and predict  $\text{CO}_2$  emissions originating from Wuhan, introducing modifications to the method that included considering chemical decay of upwind background  $\text{NO}_2$  flowing into the city. Inverse modelling approaches derived from the method of Beirle et al. (2011), as exemplified by

Lorente et al. (2019) and subsequent studies (e.g. Zhang et al. (2022), Goldberg et al. (2022), de Foy and Schauer (2022),  
60 Lange et al. (2022), Liu et al. (2022), Rey-Pommier et al. (2022)), inherently simplify the effects of atmospheric dynamics and  
chemistry. These methods have nonetheless been evaluated using synthetic data, with studies such as de Foy et al. (2014) and  
Liu et al. (2022) showing that inferred NO<sub>x</sub> emissions and lifetimes remain broadly consistent with known model input. In  
a complementary approach, Zhu et al. (2022) inferred long-term changes in NO<sub>x</sub> lifetime from decadal OMI NO<sub>2</sub> columns,  
65 using machine learning to relate NO<sub>2</sub> columns urban OH concentrations. Simplifications arise from assumptions that spatially  
and temporally varying wind speeds, NO<sub>x</sub>/NO<sub>2</sub> ratios, and NO<sub>x</sub> lifetimes may be taken as constant throughout the inversion  
domain, whereas in reality there may be substantial temporal and spatial fluctuations in these parameters, especially near the  
edges of plumes (Hakkarainen et al., 2024; Krol et al., 2024; Meier et al., 2024; Valin et al., 2013; Vilà-Guerau de Arellano  
et al., 2004). We therefore address the following research questions:

1. To what extent is the forward superposition model capable of simulating realistic NO<sub>2</sub> concentrations, despite simplifications on domain-average wind speed, NO<sub>x</sub>/NO<sub>2</sub> ratios, and NO<sub>x</sub> lifetimes?

We revisit the methodology introduced by Lorente et al. (2019) and perform an Observing System Simulation Experiment (OSSE). We generate synthetic satellite NO<sub>2</sub> observations using two Large Eddy Simulation (LES) experiments for a hypothetical city to investigate realistic chemical variations that occur in urban plumes. We assess to what extent the simple column model of Lorente et al. (2019) is capable of appropriately capturing NO<sub>2</sub> increases along with the wind over a city despite these  
75 simplifications. Next, we move on to the inversion of NO<sub>x</sub> emissions, and pose the question:

2. Can a Bayesian inversion method that weighs prior information, forward model uncertainty, and observational uncertainties improve estimates of NO<sub>x</sub> emissions from TROPOMI NO<sub>2</sub> plumes relative to a method that does not account for constraints imposed by prior knowledge?

We propose and evaluate a new more formal Bayesian inversion method, incorporating prior knowledge on NO<sub>x</sub> emissions and NO<sub>x</sub> lifetime and observations of NO<sub>2</sub> and their uncertainties. In Sections 2 and 3, we use the synthetic observations sampled from simulations with a high-resolution computational fluid dynamics model that resolves large-scale turbulence and atmospheric chemistry (MicroHH (Van Heerwaarden et al., 2017)) to evaluate the forward superposition model and inferred emissions. Then in Section 4, to demonstrate the applicability of this method, we infer a 5-year timeseries of NO<sub>x</sub> emissions for Paris on individual clear-sky days between June 2018 - August 2023 using improved TROPOMI NO<sub>2</sub>  
80 V2.4.0 retrievals. This retrieval product is based on high-resolution (0.125°) surface albedo information from the DLER database (Tilstra et al., 2023) and high-resolution (0.1°) a priori NO<sub>2</sub> profiles from CAMS (Douros et al., 2022). TROPOMI NO<sub>2</sub> retrievals based on high-resolution input data have been shown to capture NO<sub>2</sub> gradients well (Lange et al., 2024). The inversion illustrates the potential of the method and allows us to identify trends and patterns in the NO<sub>x</sub> emissions of Paris, including seasonal and weekly emission cycles, and to assess the effectiveness of pollution reduction efforts. We conclude with  
85 an evaluation of our top-down NO<sub>x</sub> estimates with an independent bottom-up inventory of NO<sub>x</sub> emissions of Paris.

## 2 Forward superposition column model

### 2.1 Model setup

The superposition column model first presented by Lorente et al. (2019) calculates NO<sub>2</sub> columns by superimposing NO<sub>x</sub> emissions along the wind within a region. The region (e.g. a city) that is analysed is divided into line cells that encapsulate the entire source region perpendicular to the wind direction (Fig. 1, (Beirle et al., 2011)). For each cell  $i$  between, within, and beyond the city length, the contribution to the NO<sub>2</sub> line density downwind to the cell is calculated using a simple column model:

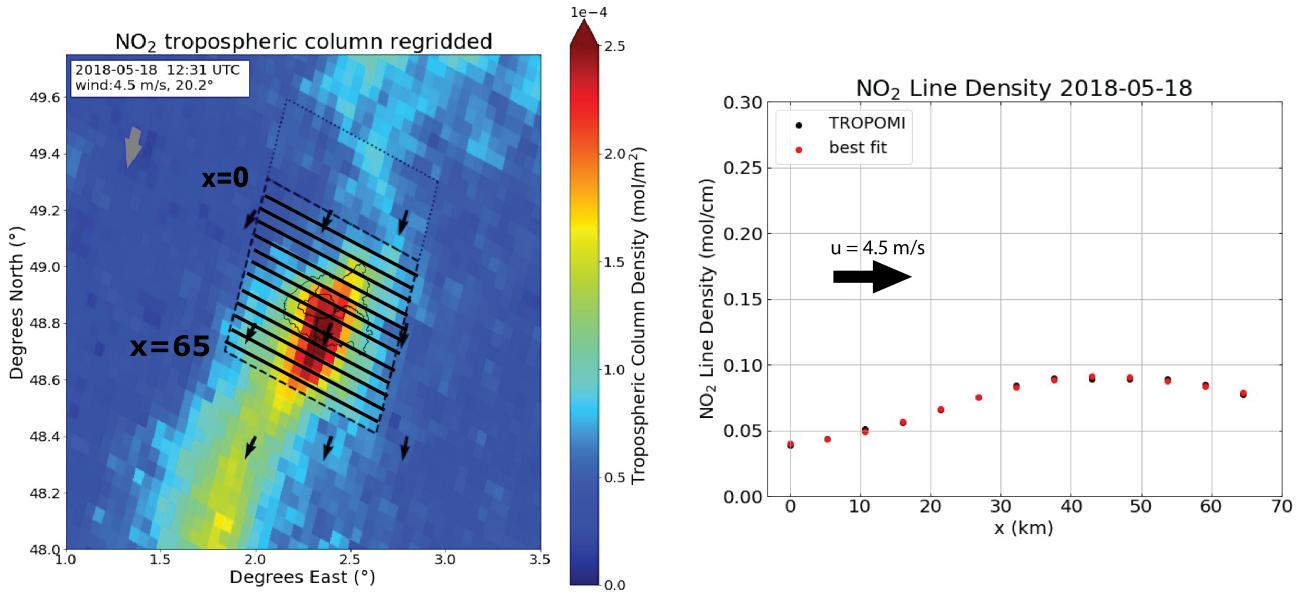
$$\begin{aligned}
 f_i(x) &= 0 && \text{for } x < x_i \\
 f_i(x) &= \frac{E_i}{k} (1 - e^{-k(x-x_i)/u}) \times \frac{[NO_2]}{[NO_x]} && \text{for } x = x_i \\
 f_i(x) &= \frac{E_i}{k} (1 - e^{-kL/u}) \times e^{-k(x-x_i)/u} \times \frac{[NO_2]}{[NO_x]} && \text{for } x > x_i
 \end{aligned} \tag{1}$$

Where  $f_i$  is the contribution of the emissions in cell  $i$  to the NO<sub>2</sub> line density at  $x$  (mol cm<sup>-1</sup>),  $E(x_i)$  represents the NO<sub>x</sub> emissions from cell  $x_i$  (mol cm<sup>-1</sup> s<sup>-1</sup>),  $L$  is the length of each line cell (m) and  $u$  is the effective wind speed at which NO<sub>x</sub> is transported (m s<sup>-1</sup>). This effective wind speed is determined by weighing the vertical wind speed profile by the the vertical NO<sub>2</sub> density profile, as described in Lorente et al. (2019). The scaling with the [NO<sub>2</sub>]/[NO<sub>x</sub>] ratio is required because a fraction of NO<sub>x</sub> is present as NO<sub>2</sub>, and TROPOMI measures the NO<sub>2</sub> columns.  $k$  is the rate constant of the chemical loss of NO<sub>x</sub> during daytime (s<sup>-1</sup>):  $k = \frac{k' [OH]}{[NO_x]/[NO_2]}$ , using the reaction rate constant  $k'$  of  $1.1 \cdot 10^{-11}$  cm<sup>3</sup> molecule<sup>-1</sup> s<sup>-1</sup> for the OH + NO<sub>2</sub> + M reaction at surface pressure for 298K (Burkholder et al., 2020), 288K (Burkholder et al., 2020). PAN formation is not explicitly considered in this framework, as it is a reversible NO<sub>x</sub> reservoir rather than a permanent sink: (e.g., (Fischer et al., 2014)). In the warm, VOC-limited conditions typical of central Paris (e.g., Johnson et al. (2024)), PAN decomposes rapidly and contributes little to net NO<sub>x</sub> loss. The dominant NO<sub>2</sub> sink under these conditions is oxidation to HNO<sub>3</sub>.

The contributions to the line density from each cell are added to the background NO<sub>2</sub> concentration ( $b$ ) to find the overall NO<sub>2</sub> line density at each distance  $x$  along with the wind:

$$F(x) = \sum_{i=1}^n f_i(x) + b \tag{2}$$

Following this model formulation, the NO<sub>2</sub> that accumulates over the city  $F(x)$  depends on ~~on~~ the spatial pattern of emissions  $E(x)$  within the city and is affected by the chemical loss and the wind speed over the city, as discussed extensively in the studies by Lorente et al. (2019) and Zhang et al. (2022). The superposition column model defined by Eqs. (1) and (2) implies that prior knowledge is required on oxidation chemistry (OH concentration and NO<sub>x</sub>:NO<sub>2</sub> ratio) within the urban boundary layer, and that the background  $b$  represents the spatially invariant free tropospheric NO<sub>2</sub> contribution to the line density.



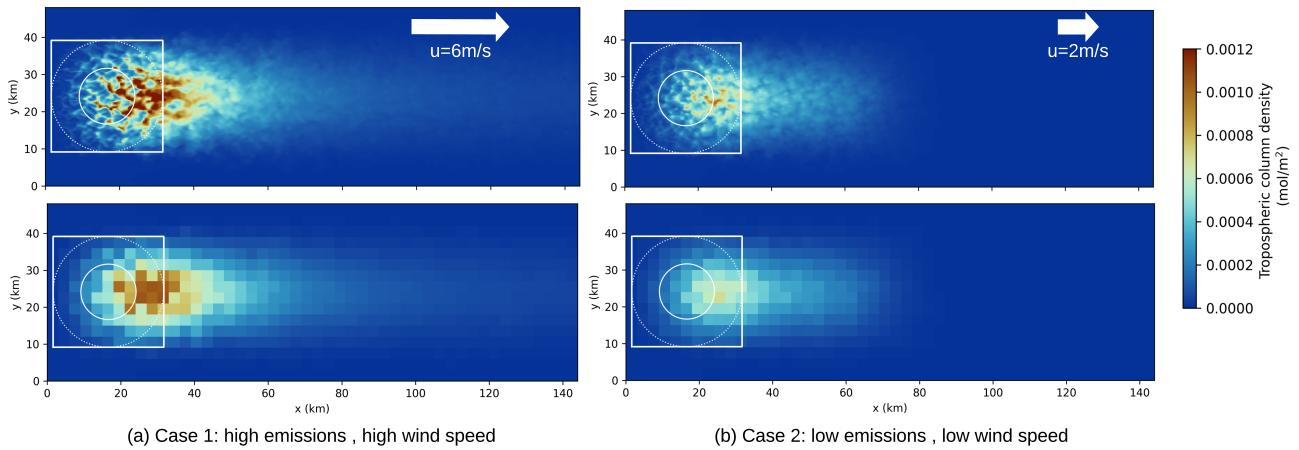
**Figure 1.** Left: the study area of wider Paris and illustration of the line density method. The black solid lines indicate the different line cells. The grey arrow indicates black arrows indicate the mean wind speed (from the northeast/north-east). The dotted grey box to the north of Paris represents the background area. Right: NO<sub>2</sub> line densities at each distance  $x$ , calculated from the TROPOMI retrieval on 18 May 2018.

## 2.2 Comparison superposition model to synthetic NO<sub>2</sub> satellite observations with MicroHH

To generate synthetic satellite observations of NO<sub>2</sub>, we use MicroHH, a direct numerical simulation (DNS) and large-eddy simulation (LES) model (Van Heerwaarden et al., 2017), which has recently been extended to include an atmospheric chemistry module based on the Kinetic Pre-Processor package (KPP) (Krol et al., 2024). We set up a horizontal model domain of 50 km (North-South)  $\times$  150 km (East-West)  $\times$  4 km (vertical), with a horizontal resolution of 300 m and a vertical resolution of 100 m. At the upper end of the model domain, a buffer zone of 750 m serves to damp gravity waves (Van Heerwaarden et al., 2017). For temperature, humidity, and momentum, circular boundary conditions were used. To avoid re-entering of emissions from the city source, we employed free outflow conditions for tracers (Ražnjević et al., 2022). More information about the MicroHH setup and initial conditions used can be found in section 1 of the Supplementary Material.

We conduct two MicroHH simulations over a hypothetical city ('symcity') to assess the capability of the forward line density model (Eq. 1,2) to realistically simulate line densities over a city, despite its simplifications (a spatio-temporally constant windspeed, NO<sub>x</sub>:NO<sub>2</sub> ratio, and NO<sub>x</sub>-lifetime). Our simulated city, 30 km  $\times$  30 km in size, is positioned on the west side of the model domain, which is dominated by a westerly flow. NO<sub>x</sub> is emitted within the city as NO, and gradually transformed to NO<sub>2</sub> by reaction with ozone. The NO emissions are spatially distributed over the city in a Gaussian pattern, so that the emissions are much larger in the center. 68% of the emissions lie within a radius of 7.5 km around the city center, and 95% within 15 km. The

simulation is performed between 6:00 and 18:00 hrs local time. Photolysis representative for the city of [Riyadh](#) is used in the simulation. We simulate two different scenario's: scenario 1 is a Spring case (photolysis of April 15th) which has high NO emissions of  $195.7 \text{ mol s}^{-1}$ , and a high wind speed of  $6 \text{ m s}^{-1}$ . Scenario 2 is a winter case (photolysis of December 15th) and has lower NO emissions of  $58.7 \text{ mol s}^{-1}$  and a wind speed of  $2 \text{ m s}^{-1}$ . Details of the two cases are summarized in Table 1. Figure 2 displays the MicroHH tropospheric  $\text{NO}_2$  columns between 0-4 km height at 13:00 hrs, close to the approximate TROPOMI overpass time of 13:30 hrs. Since the emissions are smoothly distributed over the city, the irregularities in the simulated  $\text{NO}_2$  columns, visible in the upper panels, are caused by the combined effects of atmospheric turbulence and chemistry. The lower panel shows the columns averaged to a  $5 \times 5 \text{ km}^2$  resolution, [more](#) similar to the spatial resolution of the TROPOMI  $\text{NO}_2$  retrievals.



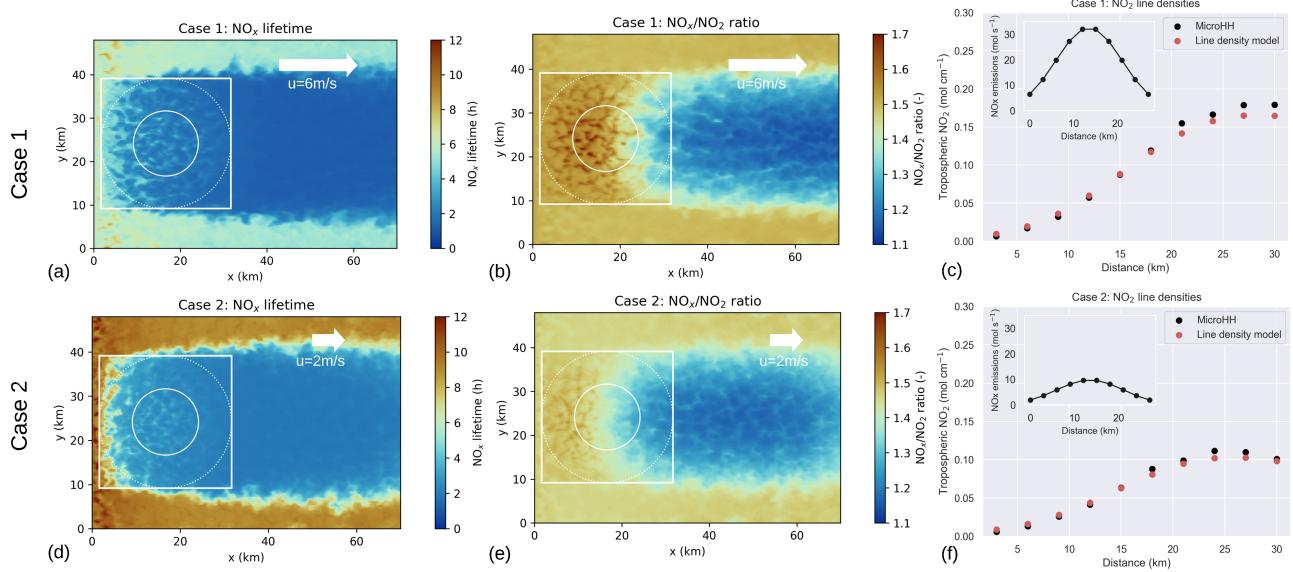
**Figure 2.** Tropospheric  $\text{NO}_2$  columns of symcity simulated by MicroHH for a Spring atmosphere with high  $\text{NO}_x$  emissions and strong wind (left panels), and a Winter atmosphere with low  $\text{NO}_x$  emissions and weak wind (right panels). 68% of the  $\text{NO}_x$  emissions occur within the smallest circle, and 95% within the bigger circle. The upper panels show the  $\text{NO}_2$  columns at the MicroHH spatial resolution of 300 m, and the lower panels show the same MicroHH simulations regridded to  $5 \times 5 \text{ km}^2$  resolution.

Building upon the  $5 \times 5 \text{ km}^2$  resampled simulations, we computed  $\text{NO}_2$  line densities across symcity, depicted as black dots in Fig. 3c and 3f. The simulated  $\text{NO}_2$  line densities over the city exhibit a tilted S-shaped pattern, similar as the observed line densities over Paris reported in Lorente et al. (2019). This pattern is a consequence of the dynamic interplay between wind and the Gaussian emission distribution across the city, with maximal emissions concentrated at the city center. NO<sub>2</sub> columns are very low upwind of the city, and the NO<sub>2</sub> line densities increase once the emitted NO is converted into NO<sub>2</sub> via the NO+O<sub>3</sub> reaction in the boundary layer. For the Spring case, with  $6 \text{ m s}^{-1}$  wind speed, the NO<sub>2</sub> line density peaks downwind of the city, reflecting the rapid transport of NO<sub>2</sub> beyond the city limits. The Winter case shows the NO<sub>2</sub> line density peaks over the city, at about 25 km, reflecting the lower wind speed in that simulation.

We evaluate the forward superposition model (based on Eq. (1) and Eq. (2)) by comparison to line densities we directly obtained from MicroHH. For the forward model, we use the NO<sub>x</sub> emissions (magnitude and spatial distribution) and other model input parameters (the average NO<sub>x</sub> lifetime, wind speed, and the average NO<sub>x</sub>/NO<sub>2</sub> ratio (Eq. 1)) from the symcity test (these average values are listed in Table 1). We try to determine these parameters as closely as possible to data we would obtain from CAMS (Copernicus Atmosphere Monitoring Service), to ensure a realistic representation of conditions encountered in a TROPOMI inversion scenario (Lorente et al., 2019). The high resolution of the MicroHH model enables us to discern the implications of these simplifications. We calculate the parameter values from the MicroHH output of 12:00 hrs local time, as the observed NO<sub>2</sub> columns ~~observed~~ depend on the conditions of preceding time steps. In the MicroHH simulation, the NO<sub>x</sub> lifetime ranges between 1-6 hours (2-9 h for case 2) within the city domain (Fig. 3a,d). The NO<sub>x</sub> lifetime is especially short in the downwind part of the plume, reflecting high OH concentrations in the urban plume (there were also substantial hydrocarbon emissions from the city, which leads to O<sub>3</sub> formation and consequently enhanced OH). The NO<sub>x</sub>/NO<sub>2</sub> ratio ranges between 1.2-1.7 (1.2-1.6 for case 2) over the city domain (Fig. 3b,e). For the Winter case (case 2) O<sub>3</sub> and OH concentrations are slightly lower than for the Spring case (case 1). This leads to a slightly longer NO<sub>x</sub> lifetime, and a lower NO<sub>x</sub>/NO<sub>2</sub> ratio than for the Spring case. In the MicroHH simulations, the wind speed remains relatively constant across the entire domain. The domain average wind speed, weighted vertically by the vertical NO<sub>2</sub> concentration is around 6 m/s for Spring and 2 m/s for Winter. Our superposition model requires effective values of wind, NO<sub>x</sub>/NO<sub>2</sub> and NO<sub>x</sub> lifetime as input parameters. Simulating NO<sub>2</sub> line densities over the city using the superposition model with the domain averaged input parameters from MicroHH yields the values shown as red dots in Fig. 3c,f. We see that despite simplifications, the simulated line densities from the superposition model closely match those from the MicroHH simulation. The agreement between the superposition and MicroHH NO<sub>2</sub> line densities allows us to estimate the superposition forward model error. We estimated the forward model error as the average absolute deviation for the 10 line density values along with the wind, ~~and~~. It amounts to 6.5% from the average MicroHH line density for case 1 and 6.3% for case 2. This suggests that, despite simplifications, the superposition model is effective and provides realistic NO<sub>2</sub> line densities based on city-domain averaged NO<sub>x</sub> emissions, lifetime and NO<sub>x</sub>/NO<sub>2</sub> ratio and wind speeds, at least at the spatial resolution of TROPOMI NO<sub>2</sub> observations.

**Table 1.** The range of parameter values within the city domain for the two cases of the MicroHH simulations for a Riyadh-like city, at 13:00, and the domain-averaged parameters used in the forward model simulations. The wind speed used in the forward model is the height-averaged wind speed weighted by the vertical NO<sub>2</sub> distribution. The range displayed for MicroHH is the city-domain averaged wind speed at 100 m height and 3 km height.

	Photolysis		Total E <sub>NO<sub>x</sub></sub> <del>Total E<sub>NO<sub>x</sub></sub></del> (mol s <sup>-1</sup> )	Wind speed (ms <sup>-1</sup> )	NO <sub>x</sub> lifetime (h)	NO <sub>x</sub> /NO <sub>2</sub> ratio	Line densities
Case 1	15 Apr	MicroHH Forward model	195.7	4 (100 m) - 8 (3 km) 5.7	<del>16.2.6±1.2</del> <del>1.7</del> 2.21	<del>1.5±0.08</del> 1.48	Black dots in Fig. Red dots in Fig.
Case 2	15 Dec	MicroHH Forward model	58.7	2 (100 m) - 3 (3 km) 2.3	<del>2.9.3.4±1.8</del> 2.87	<del>1.2.1.6</del> <del>1.4±0.08</del> 1.4	Black dots in Fig. Red dots in Fig.



**Figure 3.**  $\text{NO}_x$  lifetimes and  $\text{NO}_x/\text{NO}_2$  ratios for the two cases as simulated by MicroHH over the entire domain. Simulated line densities over ~~simcity~~ ~~symcity~~ of MicroHH and the line density model (c,f). The emission profile over the city is displayed in the small panel. The city-domain averaged conditions are given in the white box. Figures are for 13:00 hrs local time.

### 3 Inversion of $\text{NO}_x$ emissions and lifetime

We now assess the ability of the inverse superposition model to estimate known input  $\text{NO}_x$  emissions and lifetimes based on the  
175  $\text{NO}_2$  line densities as simulated by MicroHH, again using the two MicroHH simulations regridded to a TROPOMI resolution.

#### 3.1 Inversion methods

Since TROPOMI measures  $\text{NO}_2$  columns, we need to estimate the  $\text{NO}_x$  emissions using auxiliary and a priori knowledge  
180 of wind, chemical regime and emissions. In Lorente et al. (2019) a simple inversion method is used for this: the forward  
model (Eq. 1, 2) is fitted to the observed  $\text{NO}_2$  line densities by minimizing the sum of the squares of the residuals, using a  
pre-calculated look-up table with a large number of  $\text{NO}_2$  line densities corresponding to combinations of  $\text{NO}_x$  emissions and  
185  $\text{NO}_x$  lifetimes, which are allowed to vary by  $\pm 50\%$  from their prior values. No formal weight is assigned to prior knowledge  
regarding  $\text{NO}_x$  emissions and lifetimes. The optimal solution is determined by the  $\text{NO}_x$  emissions and lifetimes that result  
in the lowest residuals between observed and pre-calculated line densities. The optimal solution (with the lowest residuals)  
according to the procedure in Lorente et al. (2019) may therefore include some estimates of  $\text{NO}_x$  emissions and lifetimes that  
190 are unrealistic, e.g. with large emissions accompanied by large OH.

Here we propose a new, more formal inversion method that uses the minimization of the Bayesian cost function, taking into account knowledge on uncertainties of the prior emissions and lifetime, uncertainties in the forward model (see section 2.2), and uncertainties in the measured line densities:

$$J(x) = \frac{(x - x_a)^2}{\sigma_A^2} + \frac{(F(x) - y)^2}{\sigma_O^2} \quad (3)$$

190 Where  $x$  is the state vector, including all the terms that are fitted ( $k$ , and the emissions from each line cell). The cost function is minimized by finding the solution of  $dJ/dx = 0$ . The cost function comprises two different terms. The first term is the deviation of the prior estimate ( $x_a$ ) of the state from the actual state ( $x$ ). The second term is the deviation of the calculated line densities ( $F(x)$ ) from the measured line densities ( $y$ ), given the solution for the state  $x$ . Both terms are weighted by their uncertainties.  $\sigma_A$  represents the uncertainty in the prior, and  $\sigma_O$  represents the combined measurement uncertainty and the uncertainty in the model representation of the system. With this Bayesian inversion method, we take into account not just the observations but also our prior knowledge. This prevents the model from excessively conforming to the observed data ("overfitting"), which is problematic in the basic inversion method. The main differences between the least-squares inversion method from Lorente et al. (2019) and the Bayesian inversion approach are displayed in Table 2.

**Table 2.** Differences between the least-squares and the Bayesian inversion method in inferring the  $\text{NO}_x$  emissions and lifetimes from the  $\text{NO}_2$  line densities.

	(1) least-squares inversion	(2) Bayesian inversion
Cost function	$J(x) = (F(x) - y)^2$ (4)	$J(x) = \frac{(x - x_a)^2}{\sigma_A^2} + \frac{(F(x) - y)^2}{\sigma_O^2}$ (5)
Condition state parameters	$k\tau \pm 0.5k\tau$ $E \pm \infty$	$k\tau \pm \sigma_{A,k}\sigma_{A,\tau}$ $E \pm \sigma_{A,E}$

### 3.2 Symcity emission inversion

200 We now apply the two inversion methods to infer the  $\text{NO}_x$  emissions and lifetimes of symcity based on the  $\text{NO}_2$  columns from MicroHH, for both the high and low emission scenario. First, we assume zero observational error, and only uncertainty in the model representation of the system. We use a  $\sigma_O$  of 6%, representing the model representation uncertainty that we determined in section 2.2. We use a prior lifetime uncertainty  $\sigma_{A,k}\sigma_{A,\tau}$  of 30%, and a prior emission uncertainty  $\sigma_{A,E}$  of 50%. First, we assume known prior conditions, so the emission profiles over the city and city-average lifetimes from MicroHH 205 are used as the prior for the Bayesian model. The results of the inversions are presented in Table 3. The Bayesian inversion method yields emissions that closely match those input into MicroHH (within 2%). The least-squares inversion method slightly underestimates emissions (4%) and overestimates lifetimes for case 1, and vice versa for case 2 (16% emission overestimation)

highlighting overfitting issues inherent in the least-squares inversion approach. Although the Root Mean Square Error (RMSE) is smaller when using the least-squares inversion method, there are slight discrepancies in the inferred emissions and lifetimes  
210 due to this overfitting.

**Table 3.** Comparison of  $E_{\text{NO}_x, \text{tot}}$ ,  $\tau_{\text{NO}_x}$ , and RMSE values for different inversion methods: MicroHH, least-squares inversion, and Bayesian inversion.

	Case 1			Case 2		
	$E_{\text{NO}_x, \text{tot}}$ (mol s <sup>-1</sup> )	$\tau_{\text{NO}_x}$ (h)	RMSE (mol s <sup>-1</sup> )	$E_{\text{NO}_x, \text{tot}}$ (mol s <sup>-1</sup> )	$\tau_{\text{NO}_x}$ (h)	RMSE (mol s <sup>-1</sup> )
MicroHH	195.7	2.21		58.7	2.87	
Least-squares inversion	188	3.31	0.002	69.9	2.26	0.0016
Bayesian inversion	196.7	2.77	0.003	57.9	3.09	0.0021

This first inversion was performed for the idealized scenario, with zero observational uncertainty. To enhance the similarity of the NO<sub>2</sub> columns in MicroHH to what TROPOMI would observe, we introduced uncertainty on top of the MicroHH-simulated NO<sub>2</sub> columns. We prescribe this uncertainty as:

$$\sigma = 0.4 \cdot 10^{15} + 0.2 \cdot N_v \text{ (molecules cm}^{-2}\text{)} \quad (6)$$

215 Where  $N_v$  is the NO<sub>2</sub> column. The first part represents random uncertainty in TROPOMI measurements (originating from measurement noise in the satellite level-1 data), while the second part accounts for systematic uncertainty in estimating the Air Mass Factor (AMF), for instance caused by uncertainties in albedo estimates (Van Geffen et al., 2022; Riess et al., 2023). ~~This uncertainty is partially~~ ~~The systematic part is~~ correlated between adjacent cells (e.g. Rijssdijk et al. (2024), with a Gaussian-like shape between adjacent cells with a spatial correlation length (where the correlation falls to 1/e) of 7km (e.g. Rijssdijk et al. (2025)). We perform 1000 inversions with different random uncertainty, drawing them from a normal distribution  
220 with the standard deviation defined in Eq. 6.

The results are displayed in Fig. 4. Here we show the mean emission estimates for all 1000 runs with uncertainty on the NO<sub>2</sub> columns, for the true MicroHH emissions, the prior estimate, the Bayesian emission inversion and the Least-squares inversion.  
225 We do this for both case 1, the Spring case with high emissions (Fig. 4a) and case 2, the Winter case with lower emissions (Fig. 4b). In this more realistic setup, which uses realistic TROPOMI uncertainties instead of the idealized inversion of Table 3, emissions are more frequently overestimated in both MicroHH cases when using the least-squares inversion method than when using the Bayesian method.

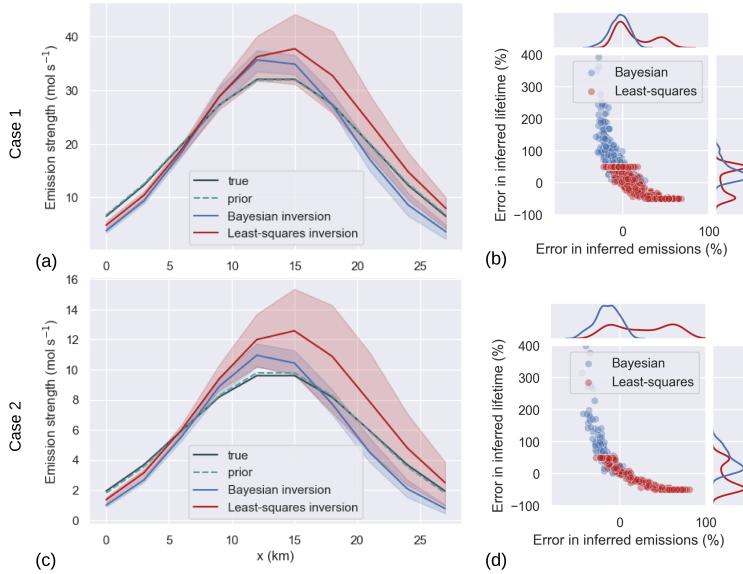
We also investigate the extent to which NO<sub>x</sub> emissions and lifetime can be independently and simultaneously reproduced.  
230 Figure 4e<sub>b,d</sub> shows the correlation between the errors in the inferred lifetime and emissions. The observed errors in the inferred NO<sub>x</sub> emissions and lifetimes exhibit substantial correlation. An overestimation in emissions is consistently accompanied by an underestimation in lifetime and vice versa. This relationship is in line with the fundamental equation of the superposition column model (Eq. 2), where both elevated emissions and extended lifetimes contribute to increased line density values. These

inversions of the emissions from MicroHH provide evidence and quantitative insights into the strength of this correlation.

235 The figures demonstrate a nonlinear relationship between the errors in inferred NO<sub>x</sub> emissions and lifetimes, where the cross-correlation appears strongest when the lifetime is underestimated. This points to a logarithmic cross-correlation between the error in NO<sub>x</sub> lifetimes and emissions, suggesting that small deviations in lifetime have a more substantial impact on emissions when the lifetime is underestimated.

240 The deviations in the inferred emissions are generally larger for the least-squares inversion method. Figures 4b,d clearly show that the lifetimes tend to be estimated at either the -50% or +50% limit that is restricted in the fit. Because of this, the emissions are also estimated in two modes, where one is close to the true value and in the other case, emissions are overestimated. For the Bayesian model, the fit is kept in check by the prior estimate, preventing overfitting, and leading more often to the right emission estimate. For the Bayesian inversion, the median error in the emission estimate of case 1 (case 2 in brackets) is -0.7% (-5.1%) and the standard deviation is 6.9% (11.3%). For the least-squares inversion, the median is 14% (26%) and the standard deviation is 22% (34%). Regarding the lifetime, the median error in the lifetime estimate of case 1 (case 2 in brackets) is 32% (13%) for the Bayesian inversion and the standard deviation is 31% (61%). For the least-squares inversion, the median is 2% (-19%) and the standard deviation is 46% (37%). To investigate the sensitivity of our results to deviations in the prior, we conducted an additional test. We repeated the OSSEs for both Symcity cases 50 times, introducing a  $\pm 20\%$  deviation in either the prior lifetime or emissions. These sensitivity tests show that increasing or decreasing the prior emissions by 20% results in a posterior bias of no more than 6% compared to the case with a known prior. Detailed results can be found in section 2 of the Supplementary Material.

245 We showed here that the strong correlation between errors in estimated NO<sub>x</sub> emissions and lifetimes makes their independent estimation difficult. The Bayesian inversion outperforms the least-squares method in estimating NO<sub>x</sub> emissions with more accurate and consistent results while avoiding the bimodal errors of the least-squares approach. Lifetimes, however, remain more challenging to reproduce and show mixed results between the two methods. The sensitivity tests show that also with deviating prior information, the Bayesian inversion method outperforms the Least-Squares approach, producing smaller biases and a smaller standard deviation.



**Figure 4.** a,c) Mean emission strength of each 5 km line cell along the city for an ensemble of 1000 randomly generated noise profiles over the city for both cases. The shaded areas indicate the interquartile range. b,d) The correlation between the deviation from the true  $\text{NO}_x$  emissions and lifetimes, with kde (kernel density estimate) plots in the margins. For the Bayesian inversion (blue) and the least-squares inversion (red).

#### 4 Model application: $\text{NO}_x$ emission estimations Paris

We now apply the two inversion methods of the superposition column model (the least-squares method and the Bayesian method) to daily clear-sky TROPOMI  $\text{NO}_2$  data to estimate  $\text{NO}_x$  emissions of Paris between June 2018 - July 2023 at the TROPOMI overpass time of around 13:00 local time. We use the [improved TROPOMI V2.4.0 data product](#), which [European TROPOMI  \$\text{NO}\_2\$  product](#) that uses CAMS a priori  $\text{NO}_2$  profiles in the air mass factor and averaging kernel calculation ([Douros et al., 2022](#)). This product is based on the operational TROPOMI  $\text{NO}_2$  (v2.4.0) version, and is described in [Douros et al. \(2022\)](#). We do not apply averaging kernels (Eskes and Boersma, 2003), because the superposition column model does not provide tropospheric  $\text{NO}_2$  profiles. Comparisons with independent  $\text{NO}_2$  data revealed an improvement in the overall bias of The standard and European  $\text{NO}_2$  product has been compared with ground-based remote sensing measurements of nine Multi-AXis Differential Optical Absorption Spectroscopy (MAX-DOAS) instruments by [Douros et al. \(2022\)](#). They found an average bias over all stations of the standard TROPOMI version of -31%. For the European product, this bias is -19%. As compared to the standard S5P tropospheric  $\text{NO}_2$  when using this TROPOMI product column data, the overall bias of the European product for almost all stations is 5 % to 18% smaller, with  $\text{NO}_2$  columns up to 30% higher than previous in emission hotspots, especially in Winter ([Douros et al., 2022](#)), providing. This provides good confidence in using TROPOMI tropospheric  $\text{NO}_2$  columns from the European product for the purpose of estimating  $\text{NO}_x$  emissions.

## 4.1 Inferring NO<sub>x</sub> emissions from TROPOMI NO<sub>2</sub> columns

### 4.0.1 Computation TROPOMI line densities

275 The TROPOMI data are first re-gridded and rotated in the wind direction. A quality check is applied where less than 10% of the TROPOMI pixels in the study area are allowed to be below a quality assurance (QA) value of 0.75 (recommended by van Geffen et al. (2022)). In this way, cloudy days or problematic retrievals are filtered out.

## 4.1 Inferring NO<sub>x</sub> emissions from TROPOMI NO<sub>2</sub> columns

### 280 4.1.1 Computation of TROPOMI line densities

For the calculation of the line densities, the TROPOMI NO<sub>2</sub> data is first rotated towards the effective wind direction (elaborated in the next section) and re-scaled into grid cells of  $0.05 \times 0.05^\circ$ . Specifically, we do this by generating a target grid with a  $0.05^\circ \times 0.05^\circ$  resolution, aligned parallel to the wind direction at the time of the TROPOMI overpass. The TROPOMI NO<sub>2</sub> data are then regredded onto this new grid, using weights based on the overlapping areas between the original and target grids.

285 Within the  $65 \times 65$  km study domain, the grid cells are divided into 13 'line cells' along the wind direction, as illustrated in Fig. 1 for the TROPOMI overpass on 18 May 2018. Subsequently, the line densities are calculated by accumulating all the pixels within each line cell, and dividing by the total width. The result is one value of the NO<sub>2</sub> line density for each 'line cell', with units of mol/cm. This is a transformation of a 2-D ( $65 \times 65$  km<sup>2</sup>) field into a 1-D line density, which simplifies the analysis, at the cost of giving up any constraints on the across-wind emission distribution.

290 The result of this line density transformation is shown for the overpass on 18 May 2018 in the right panel of Fig. 1. Here, the line densities are shown for each distance, where  $x = 0$  is the upwind start of the area. In the case of this example, the line densities are increasing until 40 km, followed by a slight decay. This pattern arises from higher emissions in the center of Paris, at  $x \sim 30$  km, after which emissions are lower and decay of NO<sub>2</sub> dominates (Lorente et al., 2019).

### 4.1.2 Estimating NO<sub>x</sub> emissions and lifetimes

295 To compute the NO<sub>x</sub> lifetimes and emissions across the city, the superpositon model (Eq. 1) is fitted to the calculated TROPOMI NO<sub>2</sub> line densities using the two different inversion methods: the least-squares method and the Bayesian variant.

The background NO<sub>2</sub> level ( $b$ ) is defined as the average line density value in a box of 30 km upwind of the study area (light grey dotted box in Fig. 1). In this definition, chemical loss of background NO<sub>2</sub> by reaction with OH is neglected, because the background NO<sub>2</sub> is assumed to be mostly located above the boundary layer where OH concentrations are assumed to be lower 300 than in the photochemically active boundary layer. The domain average, boundary layer mean NO<sub>x</sub>/NO<sub>2</sub> ratio over Paris is taken from CAMS ( $0.4^\circ \times 0.4^\circ$  resolution), 1h before the TROPOMI overpass time. The CAMS 0.4°forecast product is part of the Copernicus Atmosphere Monitoring Service and provides global 5-day forecasts of atmospheric composition at approximately 0.4° (40 km) resolution. Similarly, we take CAMS boundary layer mean OH as prior in our Bayesian inversion. Using CAMS

305 OH as a prior is justified by its physical consistency, full spatiotemporal coverage, and compatibility with the scale of the rotated line densities and column model, which assumes a single effective NO<sub>x</sub> lifetime. The domain and boundary layer average wind speed, weighted by the vertical NO<sub>2</sub> concentration from CAMS, is taken for 1h before the TROPOMI overpass from ERA5, the fifth-generation ECMWF (European Centre for MediumRange Weather Forecasts) atmospheric reanalysis of the global climate.

#### 4.1.3 Prior estimates and uncertainties

310 The TNO-MACC-III NO<sub>x</sub> emission inventory of 2011 (Kuenen et al., 2014) is used as a prior estimate of the NO<sub>x</sub> emissions over Paris. This inventory predicts a total prior NO<sub>x</sub> emission of 52.8 mol s<sup>-1</sup> over Paris for 2011. We scale this value of 2011 to the years 2018-2023 using predicted NO<sub>x</sub> emission reductions after 2011 of France by the EEA ranging from -27% for 2018 to -49% for 2023 (EEA, 2023). In 2020 and 2021, France took measures to prevent the spread of the coronavirus outbreak (Covid-19), which caused reductions in industrial activities and traffic intensity. To correct for this decrease in activity, we account for 315 an additional decrease in the prior emissions of 40% during the three Covid-19 lockdown periods of France (Guevara et al., 2021; Ikhlasse et al., 2021). The NO<sub>2</sub> concentrations in Paris never completely ~~decreased to their original levels in between~~ increased to their pre-Covid levels in between the lockdown periods (Pazmiño et al., 2021), which is why we assume that prior emissions were reduced by 20% in between the lockdown periods. A timeseries of these prior NO<sub>x</sub> emissions is displayed in the light green line of the upper panel of Fig. 5.

320 For a more realistic prior value than one yearly average, we scale the emissions using monthly, weekly and hourly emission factors from TNO (Denier van der Gon et al., 2011), based on prior knowledge of human activities. These factors are given per source sector. We weigh the temporal emission factors by the contribution of the source sectors of Paris according to the European Union (EU, 2019). This gives us an hourly scaling factor of 1.17 for 12:00 (around the TROPOMI overpass time), because of higher traffic intensity at 12:00 relative to the 24-hour mean values. We also apply additional scaling factors varying 325 per month and weekday. These are higher for Winter months, when there is more residential heating, and vehicles suffer more from so-called cold starts (Tu et al., 2021) and lower for weekend days because of reduced traffic intensity. A timeseries of this corrected prior NO<sub>x</sub> emission, including the scaling factors, is displayed in the dark green line of the upper panel of Fig. 5. It should be noted that our constructed prior is intentionally simple and straightforward. The prior only needs to be a good starting point, but the observations will adjust the posterior emissions towards the most accurate solution.

330 For the prior estimate of the NO<sub>x</sub> lifetime, the domain-average, boundary layer mean hydroxyl radical (OH) concentration is taken from CAMS for one hour before the TROPOMI overpass.

In the Bayesian inversion method, NO<sub>x</sub> emissions and lifetimes are permitted to deviate from the prior, constrained by observation and prior uncertainties. We use a line density observation uncertainty  $\sigma_o$  of 10%, accounting for both the measurement uncertainty and the uncertainty of the model representation of the system. To incorporate the uncertainty in OH concentrations 335 and its impact on the NO<sub>x</sub> lifetimes, we choose a standard deviation of 30% on the prior lifetime ( $\sigma_{\text{A},\text{K}}$ ). This selection aligns with the typical range of uncertainty observed in NO<sub>x</sub> lifetimes, which commonly falls within 50% (Lorente et al., 2019). By adopting a standard deviation of 30%, we encompass the majority of uncertainties within the expected 50% range, while

also allowing for larger deviations in exceptional cases. Finally, to account for uncertainty in the TNO-MACC-III inventory, we assign a standard deviation of the prior emissions ( $\sigma_{A,E}$ ) of 30% in each individual point (Kuenen et al., 2014)).

## 340 4.2 NO<sub>x</sub> emission estimations Paris

### 4.2.1 Emission estimates 2018-2023

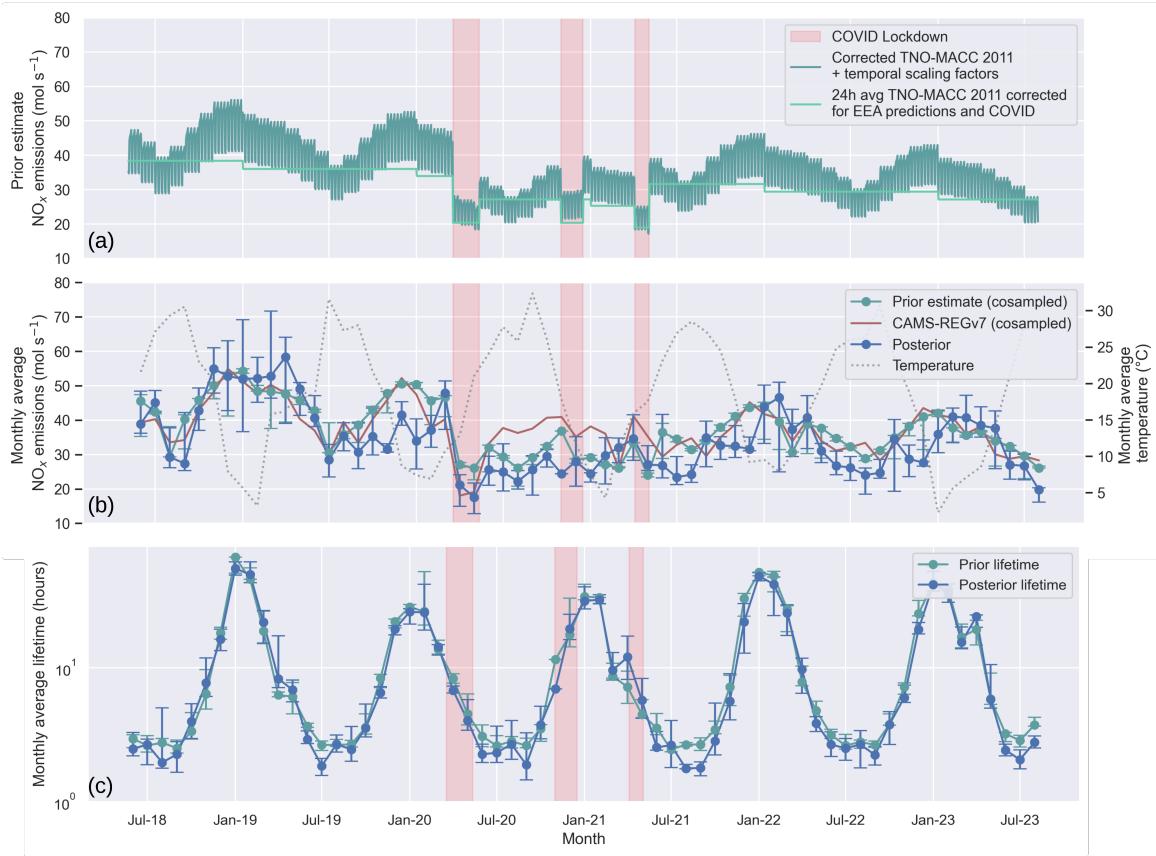
For the period spanning May 2018 to July 2023, we obtained 752 TROPOMI NO<sub>2</sub> retrievals over Paris that are under predominantly clear-sky conditions and of sufficient quality to perform NO<sub>x</sub> emission estimates. Some of these estimates are in duplicate because Paris is observed from two subsequent satellite overpasses on some days. This results in 560 inversions corresponding to unique days. Initially, the emission inversions were conducted with both the Bayesian method and the method outlined by Lorente et al. (2019). The time-averaged estimates of both methods were similar, but the latter revealed substantial outliers in NO<sub>x</sub> emission estimates, reaching up to 250 mol s<sup>-1</sup> on some days, and lifetimes were underestimated, particularly on days with low wind speed. This underscores the overfitting issue that we raised in the previous section. We therefore continue analysing the NO<sub>x</sub> emission estimates from the Bayesian inversion. Table 4 shows the average conditions across all inversions.

**Table 4.** Average meteorological and chemical conditions over Paris throughout the period May 2018 - July 2023, and prior and posterior NO<sub>x</sub> emissions and lifetimes during the TROPOMI overpass time (around 12:00). NO<sub>x</sub>/NO<sub>2</sub> ratios and wind speeds are derived from CAMS, and temperatures are measurements from the Montsouris weather station in Paris. Averages are given for the whole period and for the Summer and Winter months separately, the range is one standard deviation. The posterior standard deviation is estimated through a Monte Carlo analysis, using ~~100~~<sup>50</sup> randomly drawn prior and observation values, with their prescribed uncertainties as standard deviations, for ~~one specific day (16/06/2023)~~<sup>eight distributed days in 2022</sup>.

	n	NO <sub>x</sub> /NO <sub>2</sub>	Wind speed (m s <sup>-1</sup> )	Temperature (°C)	NO <sub>x</sub> lifetime (h)		NO <sub>x</sub> emissions (mol s <sup>-1</sup> )	
					Prior	Posterior	Prior	Posterior
Year-round	752	1.40 ± 0.12	4.6 ± 2.3	20.4 ± 8.1	5.2 ± 30%	6.0 ± <del>143.9</del> <sup>9</sup> %	35 ± 30%	32 ± <del>106.1</del> <sup>1</sup> %
Winter (DJF)	114	1.53 ± 0.19	5.4 ± 2.8	9.5 ± 4.4	30	27	39	40
Summer (JJA)	243	1.33 ± 0.09	4.0 ± 1.9	28.4 ± 3.8	2.7	2.3	31	28

350 The CAMS-derived domain average NO<sub>x</sub>/NO<sub>2</sub> ratios, averaging to 1.4 over all inversions, exceeded the commonly adopted ratio of 1.32. This discrepancy results in higher NO<sub>x</sub> emission estimates than if the constant value of 1.32 would be used. Daily temperatures, recorded at 13:00 from the Montsouris weather station in Paris city center, consistently appear relatively high, potentially influenced by urban heat island effects and our clear-sky sampling. Our findings indicate slightly lower average NO<sub>x</sub> emissions than the prior estimates, especially during the Summer months. We find an average top-down NO<sub>x</sub> emission over Paris of 32 mol s<sup>-1</sup>, which is slightly lower than the prior (9% year round), especially during the Summer months (11%).

The ~~lower panel~~<sup>middle panel (b)</sup> of Fig. 5 displays the monthly average NO<sub>x</sub> emissions of Paris estimated with the Bayesian inversion method (blue line). Values from the CAMS-REGv7 inventory are displayed in red. This is an improved version of the v4 dataset described by Kuenen et al. (2022). We scale these values using emission factors from Guevara et al. (2020, 2021)



**Figure 5.** Upper panel:  $\text{NO}_x$  emissions from the TNO-MACC III emission inventory, corrected for the emission reductions predicted by the EEA, and for emission reductions during COVID (light green). The dark green line shows daily predictions, accounting for the weekly and monthly cycle and uses a scaling factor of 1.17 (relative to the 24-hour mean) for the TROPOMI overpass time. Lower [panel](#)[two panels](#): monthly median values of the prior (green) and posterior (blue)  $\text{NO}_x$  emission [\(b\)](#) and lifetime [\(c\)](#) estimates. The prior is resampled to the days with TROPOMI  $\text{NO}_2$  retrievals. The errorbars represent the interquartile range within each month.

(more information about how we calculated these values can be found in section 2.3 of the Supplementary Material). To ensure 360 a fair comparison, the prior estimate in this graph is resampled on days with valid inversions, resulting in a slight variation from the upper panel. The monthly average posterior  $\text{NO}_x$  emissions exhibit more variability than the prior, [indicating](#). [Higher variability of posterior emissions is expected because of uncertainties in their derivation. Additionally, posterior emissions reflect real day-to-day and diurnal fluctuations, while prior emissions are based on climatological averages and are therefore inherently less variable. This difference between prior and posterior  \$\text{NO}\_x\$  emissions indicates](#) that factors beyond the month 365 and day of the week influence  $\text{NO}_x$  the emissions.

We observe an overall decreasing trend from 44 to 32 mol s<sup>-1</sup> (17.527%) in NO<sub>x</sub> emissions between May 2018/19 and August 2022/23. This decreasing trend can be partly attributed to the Paris low-emission zone, which was estimated to reduce traffic NO<sub>x</sub> emissions by about 20% between 2018-2023 due to the adoption of cleaner vehicles (Bernard et al., 2020). Significantly lower NO<sub>x</sub> emissions are visible during the Covid-19 lockdown periods, even lower than our prior assumptions. Especially 370 during the first Covid-19 lockdown (17/03/2020-11/05/2020) the NO<sub>x</sub> emissions dropped substantially. We quantified the effect of the Covid-19 lockdowns by calculating the change in emissions between the lockdown periods in 2020-2021 and the prior of the same periods in 2019. We find a significant decrease in the posterior for the first lockdown to 17.5 mol s<sup>-1</sup>, a reduction of 61%, surpassing the prior prediction of 44%. The second lockdown exhibits a reduction of around 40% compared to the 2019 prior for both the prior and posterior estimates. In the last lockdown, the decrease is less intense at 37% compared 375 to the prior's 49%.

The lower panel of Fig. 5 shows the time series of prior and posterior NO<sub>x</sub> lifetime estimates over Paris. Our Bayesian inversion framework captures seasonality, but the retrieved lifetime values should not be overinterpreted as chemically precise quantities. Our end-to-end test (Table 3 and Fig. 4) showed that the lifetime retrievals are subject to significant biases—up to 380 30%—highlighting the limitations of the method. This bias stems from the inherent asymmetry in the inversion sensitivity: the NO<sub>2</sub> line density is strongly and directly influenced by the strength of the NO<sub>x</sub> emissions, whereas the lifetime exerts a more subtle control through the dampening of the increase in line densities with distance. In practice, the signal from NO<sub>x</sub> emissions dominates the inversion, while the NO<sub>x</sub> lifetime estimate is more a regularization parameter that prevents overfitting than a robust diagnostic of possible changes in atmospheric chemistry.

Figure 5 shows good agreement between prior and posterior lifetime estimates for most months, but in summertime the 385 posterior lifetimes are often significantly below the prior values. This indicates that the NO<sub>x</sub> emission reductions from 2018 to 2023 are not only supported by the satellite-observed changes in NO<sub>2</sub> column densities, but also require shorter effective lifetimes in the inversion to fit the observed spatial gradients. Taken together with the fact that posterior NO<sub>x</sub> emissions are consistently lower than prior values, this points unambiguously to a substantial reduction in NO<sub>x</sub> pollution over Paris. The simultaneous decrease in both posterior emissions and lifetimes, relative to the prior, reinforces the robustness of this 390 conclusion: the observed NO<sub>2</sub> pattern cannot be reconciled without assuming cleaner conditions than those represented by the prior.

#### 4.2.2 Seasonal and weekly cycle

In Winter, enhanced NO<sub>x</sub> emissions are expected due to engine cold starts and increased residential heating demand (Paris generates the energy for heating within the city itself). Our analysis reveals a distinct seasonal cycle of NO<sub>x</sub> emissions with a 395 Winter:Summer ratio of 1.38 (Fig. 6a). Contrastingly, in the prior estimate, the fall months show the highest emissions (1.28 compared to the Summer). This could suggest that residential heating starts later than expected in the prior. We find that NO<sub>x</sub> emissions are generally overestimated in the prior in Summer and underestimated in Winter. This is in line with the study of Lorente et al. (2019), who found that bottom-up emission inventories underestimate actual residential heating emissions in Winter months.

400 Additionally, a significant difference is found between  $\text{NO}_x$  emissions during low temperatures ( $<10^\circ\text{C}$ ) and high temperatures ( $>20^\circ\text{C}$ ) (Fig. 6b). We filtered the data ~~for weekdays, excluding by excluding weekends~~, lockdown periods and the Summer holiday period (July-August) to mitigate potential holiday effects. The resulting average posterior difference between low and high-temperature emissions is  $43\text{-}35 \text{ mol s}^{-1}$ , a difference of 23%. This represents a similar but slightly stronger difference compared to the prior ( $43\text{-}33 \text{ mol s}^{-1}$ ).

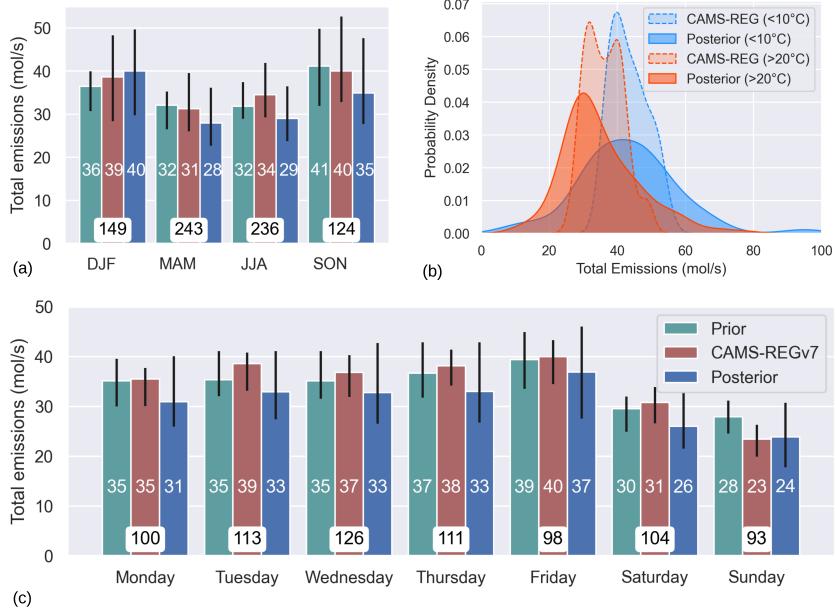
405 We observe a distinct weekly cycle in Paris (Fig. 6c), starting with low emissions on Mondays, elevated levels on Thursdays and Fridays, and a reduction during the weekend (25%). This cycle is slightly more pronounced than initially assumed in the prior, which predicts a weekend reduction, calculated relative to the Mon-Fri average, of 22%. This weekend reduction is smaller than what was found by Lange et al. (2022) (40%) and Lorente et al. (2019) (35%). In Summer, the decrease in  $\text{NO}_x$  emissions in the weekend is much larger (39%) than in Winter (11%). This is likely because of a higher contribution of 410 traffic emissions to the total emissions in the Summer months, whereas in Winter the share of traffic emissions may be smaller because of local residential heating and power generation. In Winter, the posterior weekend reduction is lower than in the prior inventory. This, again, points could point to a prior underestimation of residential heating ~~or cold start emissions in the Winter months~~. Additionally, the weaker weekly emission cycle observed in winter could be influenced by the effect of cold starts. On weekdays, vehicles are typically started early in the morning, while on weekends, car usage tends to begin later, closer to the TROPOMI overpass time. Weekend day emissions could then show up higher than without cold starts, dampening the weekly cycle.

#### 4.2.3 Literature comparison

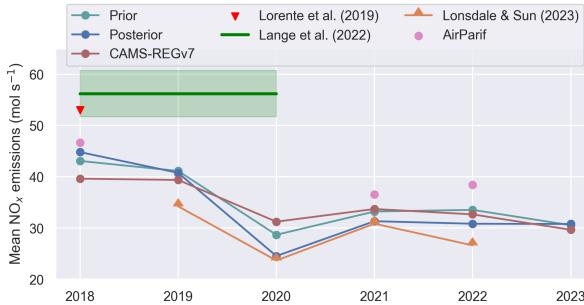
420 In Fig. 7, a comparison between prior and posterior emissions is presented alongside multiple inventory datasets. The interannual variation reveals a slight decrease in emissions in 2019, followed by a substantial decline in 2020 during the implementation of Covid-19 restrictions. Emissions veered back in 2021 and stabilized through 2023.

We compared our findings with other literature that estimates  $\text{NO}_x$  emissions in Paris (Fig. 7). Lorente et al. (2019) reported higher  $\text{NO}_x$  emissions for 2018, but only investigated these between January and June, while our estimation started from May 2018 onwards. We incorporated a corrected reaction rate constant for the oxidation of  $\text{NO}_2$ , which could contribute to the divergence in estimates. Lange et al. (2022) also estimated  $\text{NO}_x$  emissions for Paris from 2018 to 2020, reporting an average 425 emission of  $56.2 \text{ mol s}^{-1}$ . Discrepancies here may arise from a different estimation method or variations in the definition of the Paris city area. Our emission levels align more closely with those reported by Lonsdale and Sun (2023). Their findings, presented in  $\text{nmol m}^{-2} \text{ s}^{-1}$ , were converted to match our surface area unit. They observed lower values for 2019, comparable to our study for 2020 and 2021, and slightly lower values for 2022. The use of a fixed  $\text{NO}_x/\text{NO}_2$  ratio of 1.32 could contribute to their slightly lower emissions.

430 We also compare our results to the CAMS-REGv7 inventory and  $\text{NO}_x$  emission estimates from AirParif, the air quality observatory in the Île-de-France region (AirParif, 2021, 2023). AirParif provides emission estimates averaged across the entire Île-de-France area, which we scaled to align with our Paris study domain. Given that  $\text{NO}_x$  emissions are higher closer to the city center and lower in outlying areas, we applied a scaling factor of 1.7 based on the ratio of emissions from the CAMS-



**Figure 6.** The seasonal (a) and weekly cycle (c) of NO<sub>x</sub> emissions in Paris, of the prior, the posterior, and the CAMS-REGv7 inventory. White boxes display the count within each category. (c) shows the CAMS-REGv7 and Posterior NO<sub>x</sub> emissions, grouped by temperature (<10°C (n=56) and >20°C (n=189). The weekends, Summer holidays and COVID lockdowns are filtered out here.



**Figure 7.** Annual NO<sub>x</sub> emissions over the Paris from 2018 to 2023. For our study (blue) and other studies. Note that our analysis spans from May 2018 to July 2023, so the averages of 2018 and 2023 are not over the whole year.

REGv7 gridded inventory between the full Île-de-France region and our study area. These adjusted annual emission values are presented as pink dots in Fig. 7.

Since these values represent annual averages, they are not co-sampled with TROPOMI overpasses, which could explain the slightly higher NO<sub>x</sub> emission estimates from AirParif compared to our estimates (blue line). For instance, TROPOMI

overpasses with sufficient cloud-free conditions are generally more frequent in Summer, which biases our annual averages toward this season, when NO<sub>x</sub> emissions are typically lower.

440 **5 Discussion and conclusion**

We evaluated the performance of our superposition column model (Lorente et al., 2019) in estimating NO<sub>x</sub> emissions within 445 urban areas using satellite observations. Our investigation analyzed the ability of the forward model to calculate NO<sub>2</sub> columns over cities, given its use of simplified temporally and spatially averaged NO<sub>x</sub> emissions, wind speed, NO<sub>x</sub> lifetime, and NO<sub>x</sub>/NO<sub>2</sub> ratios. We performed a comparison to synthetic NO<sub>2</sub> observations generated with the high-resolution Large Eddy 450 Simulation model MicroHH, which simulates atmospheric dynamics and chemistry over a hypothetical city of 30 km × 30 km in minute detail. MicroHH simulates substantial variability in NO<sub>x</sub> lifetime (1-9 h) and NO<sub>x</sub>/NO<sub>2</sub> ratio (1.2-1.7) over the city domain, but the absolute deviation between NO<sub>2</sub> line densities simulated with the superposition model and with MicroHH stayed within 7%. This indicates that the superposition model is effective in describing the evolution of column NO<sub>2</sub> with 455 distance over a large city, given known average NO<sub>x</sub>/NO<sub>2</sub> and OH concentrations, despite averaging variable chemical and meteorological parameters over the city domain.

Tests with inferring NO<sub>x</sub> emissions from synthetic NO<sub>2</sub> line densities simulated by MicroHH using the superposition model showed that simply minimizing the least-squares using a look-up table approach, as was done before in Lorente et al. (2019) frequently resulted in overfitting, where the NO<sub>x</sub> lifetime is overestimated and the NO<sub>x</sub> emissions are underestimated or vice versa. We propose a more formal Bayesian approach of the inversion of the NO<sub>x</sub> emissions, which not only considers the fit 455 to the observations, but also incorporates prior information about NO<sub>x</sub> emissions and lifetime, to keep the solution in check. Although the Bayesian approach exhibits slightly larger discrepancies between the modeled and observed line densities, it yields solutions closer to the a priori known MicroHH emissions. The Bayesian approach reproduces the known NO<sub>x</sub> emissions to within 4%, whereas the least-squares minimization, which does not take into account uncertainties in the prior emissions, reproduces emissions to within 20%.

460 We applied Bayesian inversion to infer a 5-year time series of daily NO<sub>x</sub> emissions for Paris using TROPOMI NO<sub>2</sub> V2.4.0 retrievals from June 2018 to August 2023 under clear skies. Incorporating prior emission estimates from the TNO-MACC-III inventory, corrected for France's emission decrease reported by the European Environmental Agency, we found average NO<sub>x</sub> emissions of 32 mol s<sup>-1</sup>, which is 9% lower than the prior estimate. We observe an overall reduction in NO<sub>x</sub> emissions between 465 2018 and 2023 of 18.27%, compared to a reduction of 12% between 2018 and 2023 in the prior estimate. COVID-19 lockdowns led to sharp reductions of 61%, 40%, and 37% during the first, second, and third lockdown relative to emissions in the same period of the year before the COVID-19 measures. We observed a Winter:Summer emission ratio of 1.38, and significantly 470 higher NO<sub>x</sub> emissions on days with lower temperatures in Paris. We find a weekend NO<sub>x</sub> reduction of 25%, slightly more pronounced than the weekend effect of 22% in the emission inventory. We demonstrated that the improved Bayesian inversion method of the superposition model offers a reliable and efficient means to monitor daily NO<sub>x</sub> emissions and evaluate policies in urban areas.

*Code and data availability.* The TROPOMI L2 product used in this study is available through the TEMIS portal ([https://www.temis.nl/airpollution/no2col/no2\\_euro\\_tropomi\\_cams.php](https://www.temis.nl/airpollution/no2col/no2_euro_tropomi_cams.php), last access: 6 January 2025). CAMS model data were retrieved from the CAMS Atmosphere Data Store (<https://ads.atmosphere.copernicus.eu>, last access: 6 January 2026) and its predecessor hosted by Météo-France. The MicroHH code used for the calculations is available from GitHub (<https://github.com/microhh/microhh>, branch main\_kpp, last access: 6 January 2025). The inventory from TNO (TNO-MACC-III) is available on request by contacting HDG. Access to the CAMS-REG-v7 is provided through the Emissions of atmospheric Compounds of Ancillary Data (ECCAD) system. Since the ECCAD system requires a registration and login, a sample of the emission files has been made available for download directly. This sample includes data for the year 2017 and is available through <https://eccad.aeris-data.fr/essd-surf-emis-cams-reg/> (last access: 6 January 2025). Daily  $\text{NO}_x$  emission estimates will be made available through zenodo (Mols, 2025).

480 *Author contributions.* AM and KFB designed the study. AM performed the data analysis with support from KFB. AM wrote the manuscript with contributions from KFB, MK and HDG. MK performed the simulations in MicroHH. HDG provided the bottom-up inventory and guidance on the interpretation thereof

*Competing interests.* The authors declare that they have no conflict of interest.

## References

485 AirParif: Émissions de polluants atmosphériques et de gaz à effet de serre, Tech. rep., 2021.

AirParif: Émissions et consommations 2021 en Île-de-France (estimations faites en 2023), 2023.

Beirle, S., Boersma, K. F., Platt, U., Lawrence, M. G., and Wagner, T.: Megacity emissions and lifetimes of nitrogen oxides probed from space, *Science*, 333, 1737–1739, 2011.

Bernard, Y., Miller, J., Wappelhorst, S., and Braun, C.: Impacts of the Paris low-emission zone and implications for other cities, *TRUE–The Real Urban Emissions Initiative*, United Kingdom, ISBN, 1744072, 2020.

490 Boningari, T. and Smirniotis, P. G.: Impact of nitrogen oxides on the environment and human health: Mn-based materials for the NOx abatement, *Current Opinion in Chemical Engineering*, 13, 133–141, 2016.

Brioude, J., Angevine, W. M., Ahmadov, R., Kim, S.-W., Evan, S., McKeen, S. A., Hsie, E.-Y., Frost, G. J., Neuman, J. A., and Pollack, I. B.: Top-down estimate of surface flux in the Los Angeles Basin using a mesoscale inverse modeling technique: assessing anthropogenic emissions of CO, NO x and CO 2 and their impacts, *Atmospheric Chemistry and Physics*, 13, 3661–3677, 2013.

495 Burkholder, J. B., Sander, S. P., Abbatt, J. P. D., Barker, J. R., Cappa, C., Crounse, J. D., Dibble, T. S., Huie, R. E., Kolb, C. E., and Kurylo, M. J.: Chemical kinetics and photochemical data for use in atmospheric studies; evaluation number 19, Tech. rep., 2020.

Cheng, X., Hao, Z., Zang, Z., Liu, Z., Xu, X., Wang, S., Liu, Y., Hu, Y., and Ma, X.: A new inverse modeling approach for emission sources based on the DDM-3D and 3DVAR techniques: an application to air quality forecasts in the Beijing–Tianjin–Hebei region, *Atmospheric Chemistry and Physics*, 21, 13 747–13 761, 2021.

500 de Foy, B. and Schauer, J. J.: An improved understanding of NOx emissions in South Asian megacities using TROPOMI NO2 retrievals, *Environmental Research Letters*, 17, 024 006, 2022.

de Foy, B., Wilkins, J. L., Lu, Z., Streets, D. G., and Duncan, B. N.: Model evaluation of methods for estimating surface emissions and chemical lifetimes from satellite data, *Atmospheric Environment*, 98, 66–77, 2014.

505 Denier van der Gon, H., Hendriks, C., Kuenen, J., Segers, A., and Visschedijk, A.: TNO Report: Description of current temporal emission patterns and sensitivity of predicted AQ for temporal emission patterns, Tech. rep., TNO, Utrecht, 2011.

Douros, J., Eskes, H., van Geffen, J., Boersma, K. F., Compernolle, S., Pinardi, G., Blechschmidt, A.-M., Peuch, V.-H., Colette, A., and Veefkind, P.: Comparing Sentinel-5P TROPOMI NO 2 column observations with the CAMS-regional air quality ensemble, EGUSphere, pp. 1–40, 2022.

510 EEA: Air quality in Europe 2019, 2019.

EEA: Air quality in Europe 2022, 2022.

EEA: France – air pollution country fact sheet, 2023.

Erisman, J. W. and Draaijers, G. P. J.: Atmospheric deposition: in relation to acidification and eutrophication, Elsevier, 1995.

Eskes, H. J. and Boersma, K. F.: Averaging kernels for DOAS total-column satellite retrievals, *Atmospheric Chemistry and Physics*, 3, 515 1285–1291, 2003.

EU: Air quality: traffic measures could effectively reduce NO2 concentrations by 40% in Europe's cities, 2019.

Fischer, E. V., Jacob, D. J., Yantosca, R. M., Sulprizio, M. P., Millet, D. B., Mao, J., Paulot, F., Singh, H. B., Roiger, A., and Ries, L.: Atmospheric peroxyacetyl nitrate (PAN): a global budget and source attribution, *Atmospheric Chemistry and Physics*, 14, 2679–2698, 2014.

520 Goldberg, D. L., Harkey, M., de Foy, B., Judd, L., Johnson, J., Yarwood, G., and Holloway, T.: Evaluating NO x emissions and their effect on O<sub>3</sub> production in Texas using TROPOMI NO<sub>2</sub> and HCHO, *Atmospheric Chemistry and Physics*, 22, 10 875–10 900, 2022.

Guevara, M., Jorba, O., Tena, C., Denier van der Gon, H., Kuenen, J., Elguindi-Solomon, N., Darras, S., Granier, C., and Pérez García-Pando, C.: CAMS-TEMPO: global and European emission temporal profile maps for atmospheric chemistry modelling, *Earth System Science Data Discussions*, 2020, 1–60, 2020.

525 Guevara, M., Jorba, O., Soret, A., Petetin, H., Bowdalo, D., Serradell, K., Tena, C., Denier van der Gon, H., Kuenen, J., and Peuch, V.-H.: Time-resolved emission reductions for atmospheric chemistry modelling in Europe during the COVID-19 lockdowns, *Atmospheric Chemistry and Physics*, 21, 773–797, 2021.

Hakkarainen, J., Kuhlmann, G., Koene, E., Santaren, D., Meier, S., Krol, M. C., van Stratum, B. J. H., Ialongo, I., Chevallier, F., and Tamminen, J.: Analyzing nitrogen dioxide to nitrogen oxide scaling factors for data-driven satellite-based emission estimation methods: 530 A case study of Matimba/Medupi power stations in South Africa, *Atmospheric Pollution Research*, 15, 102 171, 2024.

Ikhlasse, H., Benjamin, D., Vincent, C., and Hicham, M.: Environmental impacts of pre/during and post-lockdown periods on prominent air pollutants in France, *Environment, Development and Sustainability*, 23, 14 140–14 161, 2021.

Jacob, D.: *Introduction to Atmospheric Chemistry*, Princeton University Press, 1999.

Johnson, M. S., Philip, S., Meech, S., Kumar, R., Sorek-Hamer, M., Shiga, Y. P., and Jung, J.: Insights into the long-term (2005–2021) 535 spatiotemporal evolution of summer ozone production sensitivity in the Northern Hemisphere derived with the Ozone Monitoring Instrument (OMI), *Atmospheric Chemistry and Physics*, 24, 10 363–10 384, 2024.

Krol, M., van Stratum, B., Anglou, I., and Boersma, K. F.: Evaluating NO x stack plume emissions using a high-resolution atmospheric chemistry model and satellite-derived NO<sub>2</sub> columns, *Atmospheric Chemistry and Physics*, 24, 8243–8262, 2024.

Kuenen, J., Dellaert, S., Visschedijk, A., Jalkanen, J.-P., Super, I., and Denier van der Gon, H.: CAMS-REG-v4: a state-of-the-art high- 540 resolution European emission inventory for air quality modelling, *Earth System Science Data*, 14, 491–515, 2022.

Kuenen, J. J. P., Visschedijk, A. J. H., Jozwicka, M., and Denier Van Der Gon, H. A. C.: TNO-MACC<sub>II</sub> emission inventory; a multi-year (2003–2009) consistent high-resolution European emission inventory for air quality modelling, *Atmospheric Chemistry and Physics*, 14, 10 963–10 976, 2014.

Kurokawa, J.-i., Yumimoto, K., Uno, I., and Ohara, T.: Adjoint inverse modeling of NO<sub>x</sub> emissions over eastern China using satellite 545 observations of NO<sub>2</sub> vertical column densities, *Atmospheric Environment*, 43, 1878–1887, 2009.

Lange, K., Richter, A., and Burrows, J. P.: Variability of nitrogen oxide emission fluxes and lifetimes estimated from Sentinel-5P TROPOMI observations, *Atmospheric Chemistry and Physics*, 22, 2745–2767, 2022.

Lange, K., Richter, A., Bösch, T., Zilker, B., Latsch, M., Behrens, L. K., Okafor, C. M., Bösch, H., Burrows, J. P., and Merlaud, A.: Validation 550 of GEMS tropospheric NO<sub>2</sub> columns and their diurnal variation with ground-based DOAS measurements, *Atmospheric Measurement Techniques*, 17, 6315–6344, 2024.

Liu, F., Tao, Z., Beirle, S., Joiner, J., Yoshida, Y., Smith, S. J., Knowland, K. E., and Wagner, T.: A new method for inferring city emissions and lifetimes of nitrogen oxides from high-resolution nitrogen dioxide observations: a model study, *Atmospheric Chemistry and Physics*, 22, 1333–1349, 2022.

Lonsdale, C. R. and Sun, K.: Nitrogen oxides emissions from selected cities in North America, Europe, and East Asia observed by the 555 TROPOMI before and after the COVID-19 pandemic, *Atmospheric Chemistry and Physics*, 23, 8727–8748, 2023.

Lorente, A., Boersma, K. F., Eskes, H. J., Veefkind, J. P., Van Geffen, J., De Zeeuw, M. B., Denier Van Der Gon, H. A. C., Beirle, S., and Krol, M. C.: Quantification of nitrogen oxides emissions from build-up of pollution over Paris with TROPOMI, *Scientific reports*, 9, 1–10, 2019.

560 Meier, S., Koene, E. F. M., Krol, M., Brunner, D., Damm, A., and Kuhlmann, G.: A lightweight NO 2-to-NO x conversion model for quantifying NO x emissions of point sources from NO 2 satellite observations, *Atmospheric Chemistry and Physics*, 24, 7667–7686, 2024.

Pazmiño, A., Beekmann, M., Goutail, F., Ionov, D., Bazureau, A., Nunes-Pinharanda, M., Hauchecorne, A., and Godin-Beekmann, S.: Impact of the COVID-19 pandemic related to lockdown measures on tropospheric NO 2 columns over Île-de-France, *Atmospheric Chemistry and Physics*, 21, 18 303–18 317, 2021.

565 Ražnjević, A., Van Heerwaarden, C., Van Stratum, B., Hensen, A., Velzeboer, I., Van Den Bulk, P., and Krol, M.: Interpretation of field observations of point-source methane plume using observation-driven large-eddy simulations, *Atmospheric Chemistry and Physics*, 22, 6489–6505, 2022.

Rey-Pommier, A., Chevallier, F., Ciais, P., Broquet, G., Christoudias, T., Kushta, J., Hauglustaine, D., and Sciare, J.: Quantifying NO x emissions in Egypt using TROPOMI observations, *Atmospheric Chemistry and Physics Discussions*, 2022, 1–21, 2022.

570 Riess, T. C. V. W., Boersma, K. F., Van Roy, W., de Laat, J., Dammers, E., and van Vliet, J.: To new heights by flying low: Comparison of aircraft vertical NO 2 profiles to model simulations and implications for TROPOMI NO 2 retrievals, *EGUphere*, 2023, 1–25, 2023.

Rijsdijk, P., Eskes, H., Dingemans, A., Boersma, F., Sekiya, T., Miyazaki, K., and Houweling, S.: Quantifying uncertainties of satellite NO 2 superobservations for data assimilation and model evaluation, *EGUphere*, 2024, 1–42, 2024.

575 Rijsdijk, P., Eskes, H., Dingemans, A., Boersma, K. F., Sekiya, T., Miyazaki, K., and Houweling, S.: Quantifying uncertainties in satellite NO2 superobservations for data assimilation and model evaluation, *Geoscientific Model Development*, 18, 483–509, 2025.

Seinfeld, J. H. and Pandis, S. N.: *Atmospheric chemistry and physics: from air pollution to climate change*, John Wiley & Sons, 2016.

Sierk, B., Fernandez, V., Bézy, J.-L., Meijer, Y., Durand, Y., Courrèges-Lacoste, G. B., Pachot, C., Löscher, A., Nett, H., and Minoglou, K.: The Copernicus CO2M mission for monitoring anthropogenic carbon dioxide emissions from space, in: *International Conference on Space Optics—ICSO 2020*, vol. 11852, pp. 1563–1580, SPIE, 2021.

580 Stavrakou, T., Müller, J.-F., Boersma, K. F., Van Der A, R. J., Kurokawa, J., Ohara, T., and Zhang, Q.: Key chemical NO x sink uncertainties and how they influence top-down emissions of nitrogen oxides, *Atmospheric Chemistry and Physics*, 13, 9057–9082, 2013.

Tilstra, L. G., de Graaf, M., Trees, V., Litvinov, P., Dubovik, O., and Stammes, P.: A directional surface reflectance climatology determined from TROPOMI observations, *Atmospheric Measurement Techniques Discussions*, 2023, 1–29, 2023.

585 Tu, R., Xu, J., Wang, A., Zhai, Z., and Hatzopoulou, M.: Effects of ambient temperature and cold starts on excess NOx emissions in a gasoline direct injection vehicle, *Science of the Total Environment*, 760, 143 402, 2021.

Valin, L. C., Russell, A. R., and Cohen, R. C.: Variations of OH radical in an urban plume inferred from NO2 column measurements, *Geophysical Research Letters*, 40, 1856–1860, 2013.

590 van Geffen, J., Eskes, H., Boersma, K., and Veefkind, J.: TROPOMI ATBD of the total and tropospheric NO2 data products: S5P-KNMI-L2-0005-RP, Tech. rep., 2022.

Van Geffen, J., Eskes, H., Compernolle, S., Pinardi, G., Verhoelst, T., Lambert, J.-C., Sneep, M., Ter Linden, M., Ludewig, A., and Boersma, K. F.: Sentinel-5P TROPOMI NO 2 retrieval: impact of version v2. 2 improvements and comparisons with OMI and ground-based data, *Atmospheric Measurement Techniques*, 15, 2037–2060, 2022.

Van Heerwaarden, C. C., Van Stratum, B. J. H., Heus, T., Gibbs, J. A., Fedorovich, E., and Mellado, J. P.: MicroHH 1.0: A computational  
595 fluid dynamics code for direct numerical simulation and large-eddy simulation of atmospheric boundary layer flows, *Geoscientific Model  
Development*, 10, 3145–3165, 2017.

Vilà-Guerau de Arellano, J., Dosio, A., Vinuesa, J.-F., Holtslag, A. A. M., and Galmarini, S.: The dispersion of chemically reactive species  
in the atmospheric boundary layer, *Meteorology and Atmospheric Physics*, 87, 23–38, 2004.

Visser, A. J., Boersma, K. F., Ganzeveld, L. N., and Krol, M. C.: European NO x emissions in WRF-Chem derived from OMI: impacts on  
600 summertime surface ozone, *Atmospheric Chemistry and Physics*, 19, 11 821–11 841, 2019.

Vitousek, P. M., Aber, J. D., Howarth, R. W., Likens, G. E., Matson, P. A., Schindler, D. W., Schlesinger, W. H., and Tilman, D. G.: Human  
alteration of the global nitrogen cycle: sources and consequences, *Ecological applications*, 7, 737–750, 1997.

WHO: WHO global air quality guidelines: particulate matter (PM2. 5 and PM10), ozone, nitrogen dioxide, sulfur dioxide and carbon  
monoxide, World Health Organization, 2021.

605 Yan, C., Nie, W., Vogel, A. L., Dada, L., Lehtipalo, K., Stolzenburg, D., Wagner, R., Rissanen, M. P., Xiao, M., and Ahonen, L.: Size-  
dependent influence of NOx on the growth rates of organic aerosol particles, *Science advances*, 6, eaay4945, 2020.

Zhang, Q., Boersma, K. F., Zhao, B., Eskes, H., Chen, C., Zheng, H., and Zhang, X.: Quantifying daily NO x and CO 2 emissions from  
Wuhan using satellite observations from TROPOMI and OCO-2, *Atmospheric Chemistry and Physics Discussions*, pp. 1–18, 2022.

Zhu, Q., Laughner, J. L., and Cohen, R. C.: Estimate of OH trends over one decade in North American cities, *Proceedings of the National  
610 Academy of Sciences*, 119, e2117399 119, 2022.