



1 *Manuscript for*

2 **Divergent Drivers of Aerosol Acidity: Evidence for**
3 **Shifting Regulatory Regimes in a Coastal Region**

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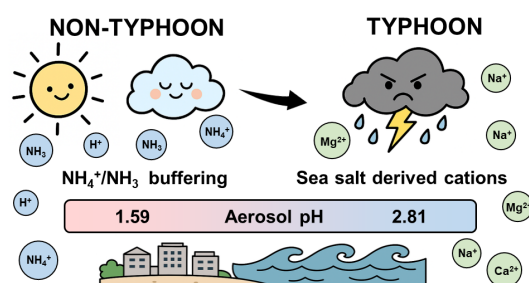
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ABSTRACT: Aerosol acidity plays a crucial role in multiphase atmospheric chemistry, influencing aerosol composition, gas-particle partitioning, and the oxidative capacity of atmosphere. However, the mechanisms governing aerosol acidity in coastal area under extreme weather remains challenging due to its complexity of atmospheric transport. Here, we investigate aerosol pH in Shenzhen, a coastal megacity in China, by integrating field observations with multiphase buffer theory and ISORROPIA simulations. Our observations captured both a typhoon episode and typical non-typhoon periods with two contrasting regimes: during non-typhoon periods, aerosols were consistently buffered by the $\text{NH}_4^+/\text{NH}_3$ pair, with relative humidity serving as the primary driver of pH variability, enabling reliable predictions using multiphase buffer theory. In contrast, during a typhoon episode, sea salt derived nonvolatile cations emerged as the dominant drivers, violating the charge balance for $\text{NH}_4^+/\text{NH}_3$ buffering and leading to poor performance of buffer theory. Under these conditions, ISORROPIA simulations with constant aerosol water content reproduced the observed pH more reliably, highlighting a compositional rather than meteorological control. Our results provide the direct field-based evidence for regime shifts in aerosol acidity regulation in coastal area, and underscore the need for chemical transport models to account for composition-meteorology interactions to improve acidity predictions under extreme weather events.

Abstract Graphic





1 INTRODUCTION

Aerosol acidity is a key regulator in multiphase atmospheric chemistry. It governs the gas-particle partitioning of semi-volatile species (e.g., NH_3 , HNO_3 , HCl , and organic acids/bases) and dictates critical aqueous-phase processes, including SO_2 oxidation, secondary transformations of organic compounds, and the activation of trace metals (Tilgner et al., 2021; Pye et al., 2020; Cai et al., 2024; Surratt et al., 2007). Through these pathways, aerosol acidity exerts strong control over atmospheric oxidative capacity and pollutant lifetimes (Pye et al., 2020). On a large scale, aerosol acidity plays a pivotal role in determining particle composition, optical properties, and hygroscopicity, thereby influencing their radiative impacts and the ability to act as cloud condensation nuclei (Turnock et al., 2019; Karydis et al., 2021; Xu et al., 2020; Zhang et al., 2023). Moreover, aerosol acidity enhances particle toxicity by directly triggering respiratory inflammation, and affects the solubility of heavy metals, thereby regulating their bioavailability in terrestrial and marine ecosystems (Fang et al., 2017; Song et al., 2024; Amdur et al., 1978; Longo et al., 2016). Variations in aerosol acidity are not only fundamental to atmospheric chemistry processes but also directly influence regional air quality management (e.g., coordinated control of $\text{PM}_{2.5}$ and ozone), the global nitrogen and sulfur cycles, and climate feedback mechanisms. Thus, advancing the understanding of aerosol acidity has critical implications for public health and environmental policy.

Precise quantification of aerosol acidity remains an important yet challenging issue in atmospheric chemistry. Aerosol acidity is typically determined by aerosol pH. Traditional methods based on filter extraction and subsequent H^+ quantification are susceptible to substantial artifacts arising from sampling and dilution (Pathak et al., 2004; Hennigan et al., 2015). Recent techniques, such as Raman-based microdroplet pH detection (Cui et al., 2021), aerosol optical tweezers (Boyer et al., 2020), fluorescence probes (Li and Kuwata, 2023), and quantitative colorimetric imaging (Craig et al., 2018), have provided novel insights into particle-scale acidity, though their applicability to real-world aerosols remains limited, particularly for real-time dynamic monitoring of aerosol acidity. Indirect proxies including ion balance and gas-to-particle molar ratios, while



63 widely applied, suffer from systematic biases owing to neglected organic acid dissociation and
64 semi-volatile partitioning (Metzger et al., 2006;Hennigan et al., 2015). Thermodynamic
65 equilibrium models have emerged as the dominant framework for estimating aerosol pH values
66 (Saxena et al., 1986;Jacobson et al., 1996;Pilinis and Seinfeld, 1987;Wexler and Seinfeld, 1991).
67 Among them, the Extended Aerosol Inorganic Model (E-AIM) and ISORROPIA are widely
68 employed (Wexler and Clegg, 2002;Nenes et al., 1998), with E-AIM regarded as the benchmark
69 owing to its explicit treatment of ion activity coefficients (Clegg et al., 1992), while ISORROPIA
70 is favored for its computational efficiency (Nenes et al., 1999). Comparative studies generally
71 report consistent pH estimates from these two models, albeit with context-dependent deviations
72 (Song et al., 2018;Hennigan et al., 2015;Battaglia et al., 2019). Substantial uncertainties persist in
73 characterizing the spatiotemporal variability of aerosol acidity and its dynamic response to
74 chemical composition and meteorological drivers. Research remains limited in coastal megacities
75 and under extreme weather events such as typhoons, where intense atmospheric transport may
76 substantially challenge the applicability of existing models and theories.

77 Aerosol acidity exhibits strong spatiotemporal variability, mainly arising from the combined
78 influences of particle chemical composition and meteorological conditions (Zhou et al.,
79 2022;Zhang et al., 2021;Ding et al., 2019;Wang et al., 2022;Tao and Murphy, 2019, 2021). In
80 particular, water-soluble inorganic components exert significant control, with sulfate substantially
81 enhancing aerosol acidity due to its low volatility, whereas nitrate, with its strong hygroscopicity,
82 increases aerosol water content (AWC) under elevated relative humidity, thereby lowering acidity
83 (Ding et al., 2019). Therefore, a lower nitrate-to-sulfate ratio generally leads to more acidic
84 particles (Xie et al., 2020). Although the direct contribution of organics to aerosol pH is relatively
85 minor (Guo et al., 2015), interactions between inorganic and organic components can alter acidity
86 with pH increases of up to 0.7 units (Pye et al., 2018). Meteorological conditions exert a strong
87 influence on aerosol acidity by altering both particle water content and gas-particle partitioning.
88 Increased relative humidity facilitates hygroscopic growth and promotes aqueous-phase reactions,
89 which typically dilute proton concentrations and thus mitigate acidity (Bian et al., 2014). In



90 contrast, higher temperatures shift the equilibrium of semi-volatile species such as ammonia
91 toward the gas phase and enhance water vapor pressure, processes that together elevate proton
92 loading and intensify aerosol acidity (Guo et al., 2018; Battaglia et al., 2017). The recently
93 proposed multiphase buffer theory provides a new framework for understanding the long-term
94 evolution of aerosol acidity by emphasizing the stabilizing role of conjugate acid-base pairs against
95 external perturbations, thereby exerting critical control over sulfate formation pathways and other
96 multiphase atmospheric processes (Zheng et al., 2020; Zheng et al., 2022; Chen et al., 2022; Gao et
97 al., 2025), highlighting the complexity of aerosol acidity driving mechanisms. However, the
98 relative importance of different buffering systems under complex meteorological conditions
99 remains poorly constrained, and the potential dominance of nonvolatile cations (NVCs, e.g., Na^+ ,
100 Ca^{2+} , Mg^{2+} , K^+) in coastal environments and during extreme weather events has not been
101 systematically assessed with field evidence.

102 In this study, we investigated the buffering capacity and controlling factors of aerosol pH in
103 Shenzhen, China, a subtropical coastal megacity frequently influenced by typhoons, by integrating
104 field observations with multiphase buffer theory and ISORROPIA II simulations. Our field
105 observations captured both a typhoon episode and typical non-typhoon periods, providing a natural
106 contrast between distinct atmospheric regimes. By comparing the buffering capacity and key
107 drivers of aerosol pH during typhoon and non-typhoon periods, we aim to shed light on the
108 complex interactions between composition and meteorology that regulate aerosol acidity. Our
109 results provide the field-based evidence the role of NVCs in modulating aerosol acidity under
110 extreme weather conditions. These findings underscore the need to account for meteorology-
111 composition interactions when applying multiphase buffer theory in coastal regions, and reveal
112 important limitations of current models under episodic perturbations.

113 2 METHODS

114 **2.1 Field Measurements.** Field measurements were carried out at the Xichong site (22.48°N,
115 114.56°E) on the Dapeng Peninsula in Shenzhen, China, from August to September 2022. Detailed



information and site characteristics of Xichong have been described elsewhere (Zhai et al., 2025). Briefly, Xichong is located at the southeastern end of Shenzhen, a representative coastal megacity in southern China. During the field campaign, a Monitor for Aerosols and Gases in Ambient air (MARGA, Metrohm-Applikon, Netherlands) was utilized to measure online concentrations of major water-soluble gases (NH_3 , SO_2 , HNO_3 , HCl) and aerosol ions (NH_4^+ , Na^+ , K^+ , Ca^{2+} , Mg^{2+} , SO_4^{2-} , NO_3^- , Cl^-). Additional measurements at the site included $\text{PM}_{2.5}$ and O_3 mass concentrations, as well as key meteorological parameters (temperature, relative humidity, wind speed, and wind direction). In this study, the analysis period used was from 24 August to 11 September 2022, corresponding to the overlapping operational time of all deployed instruments, with all online data standardized to a temporal resolution of 1 h.

2.2 ISORROPIA Calculation. The pH is defined in terms of the activity of hydrogen ions (H^+) in aqueous solution, expressed on a molality basis as:

$$\text{pH} = -\log_{10}[a(\text{H}^+)] = -\log_{10}[\chi(\text{H}^+) \cdot \gamma(\text{H}^+)] \quad (1)$$

where $a(\text{H}^+)$ is the activity of H^+ , and $\chi(\text{H}^+)$ and $\gamma(\text{H}^+)$ denote the mole fraction and mole fraction-based activity coefficients of H^+ in aqueous solution, respectively. In this study, aerosol pH is calculated using ISORROPIA II (<http://isorropia.epfl.ch>), which provides direct predictions of ionic activity, gas partial pressure, and the phase volumes of solids and liquids. Here, the model considers the $\text{Na}^+ - \text{NH}_4^+ - \text{K}^+ - \text{Ca}^{2+} - \text{Mg}^{2+} - \text{Cl}^- - \text{NO}_3^- - \text{SO}_4^{2-}$ system and is run in “forward mode” under the “metastable” phase state, which avoids salt crystallization and is more representative of ambient aerosol conditions under high humidity. The model is driven by measured concentrations of major aerosol ions and gases, along with observed meteorological parameters. Aerosol pH was then calculated from the ISORROPIA output as:

$$\text{pH} = -\log_{10}[\text{H}_{\text{aq}}^+] = -\log_{10}[1000 \cdot \text{H}_{\text{air}}^+ / \text{AWC}] \quad (2)$$

where H_{aq}^+ is the equilibrium hydronium ion concentration in ambient aerosol liquid water (mol L^{-1}), H_{air}^+ is the equilibrium hydronium ion concentration per unit volume of air ($\mu\text{g m}^{-3}$), and



141 AWC is the aerosol water content ($\mu\text{g m}^{-3}$). The factor 1000 accounts for the unit conversion
142 between $\mu\text{g m}^{-3}$ of air and mol L^{-1} of aerosol liquid water.

143 **2.3 Multiphase Buffer Theory.** A buffer system is defined as a chemical system that exhibits
144 resilience to pH perturbations upon the addition of a certain amount of acid or base. In
145 homogeneous aqueous solutions, the conjugate acid-base pairs are confined to the liquid phase,
146 and thus the solution pH is governed exclusively by acid dissociation equilibria. In multiphase
147 systems, however, the volatile components of the conjugate pairs can reside in both the gas and
148 liquid phases, whereby the pH is modulated by the coupled effects of gas-liquid partitioning and
149 dissociation equilibria. The recently proposed multiphase buffer theory introduced an analytical
150 framework for the buffer capacity β of aerosol systems (Zheng et al., 2020). Here, β is defined as
151 the amount of acid (dn_{acid}) required to reduce pH by dpH units, or the amount of base (dn_{base})
152 required to raise pH by dpH units. It thus characterizes the instantaneous buffering capacity,
153 relating infinitesimal changes in acid/base content to the corresponding change in pH. In this study,
154 β (mol kg^{-1}) is expressed as:

$$155 \quad \beta = -\frac{dn_{\text{acid}}}{dpH} = \frac{dn_{\text{base}}}{dpH} = 2.303 \left\{ \frac{K_w}{[H^+]} + [H^+] + \sum_i \frac{K_{a,i}^* [H^+]}{(K_{a,i}^* + [H^+])^2} \cdot [X_i]_{\text{tot}}^* \right\} \quad (3-1)$$

156 where K_w is the water dissociation constant ($\text{mol}^2 \text{kg}^{-2}$), $K_{a,i}^*$ represents the effective acid
157 dissociation constant of buffering agent X_i in gas-liquid multiphase systems (mol kg^{-1}), and
158 $[X_i]_{\text{tot}}^*$ is the total equivalent molality of X_i , accounting for both gaseous and aqueous phases
159 (mol kg^{-1}).

160 The self-buffering capacity of water is an inherent property conferring resistance to pH changes.
161 Here, it is defined as:

$$162 \quad \beta_{\text{water}} = 2.303 \left\{ \frac{K_w}{[H^+]} + [H^+] \right\} \quad (3-2)$$

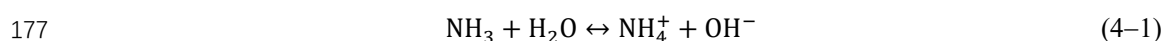
163 which becomes appreciable only under extreme acidic ($\text{pH} < 2$) or alkaline ($\text{pH} > 12$) conditions,
164 but is negligible within the typical aerosol pH range. In contrast, conjugate acid-base pairs (e.g.,
165 $\text{NH}_4^+/\text{NH}_3$, $\text{HNO}_3/\text{NO}_3^-$, $\text{HSO}_4^-/\text{SO}_4^{2-}$) dominate buffering within this range, while minor organic



acids contribute little (Zheng et al., 2020). Similar to bulk aqueous systems, the effective buffer range ($pK_{a,i}^* \pm 1$) corresponds to the pH interval over which these conjugate pairs exert substantial buffering effects.

Previous study (Zheng et al., 2020) further introduced the dimensionless gas-liquid partitioning constant (K_g) to represent the equivalent molality of gaseous species dissolved in the aqueous phase. This formulation enables explicit treatment of gas-particle equilibria for buffering agents such as $\text{NH}_4^+/\text{NH}_3$ and $\text{HNO}_3/\text{NO}_3^-$. Corrections for non-ideality due to elevated ionic strength can also be incorporated through activity coefficients. Despite simplifying assumptions, such as equilibrium thermodynamics and the neglect of organic contributions, multiphase buffer theory provides a powerful framework for quantifying aerosol buffering mechanisms.

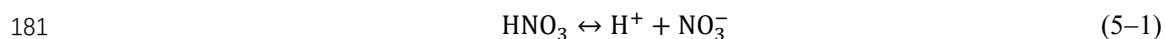
For the buffering agent $\text{NH}_4^+/\text{NH}_3$,



$$K_g = \frac{[\text{NH}_3(\text{g})]}{[\text{NH}_3(\text{aq})]} = \frac{\rho_w}{H_{\text{NH}_3} \cdot R \cdot T \cdot \text{AWC}} \quad (4-2)$$

$$K_{a,\text{NH}_3}^* = \frac{[\text{H}^+(\text{aq})] \cdot ([\text{NH}_3(\text{aq})] + [\text{NH}_3(\text{g})])}{[\text{NH}_4^+(\text{aq})]} = K_{a,\text{NH}_3} \cdot \left(1 + \frac{\rho_w}{H_{\text{NH}_3} \cdot R \cdot T \cdot \text{AWC}}\right) \quad (4-3)$$

And for the buffering agent $\text{HNO}_3/\text{NO}_3^-$,



$$K_g = \frac{[\text{HNO}_3(\text{g})]}{[\text{HNO}_3(\text{aq})]} = \frac{\rho_w}{H_{\text{HNO}_3} \cdot R \cdot T \cdot \text{AWC}} \quad (5-2)$$

$$K_{a,\text{HNO}_3}^* = \frac{[\text{H}^+(\text{aq})] \cdot ([\text{NO}_3^-(\text{aq})] + [\text{HNO}_3(\text{g})])}{[\text{HNO}_3(\text{aq}) + \text{HNO}_3(\text{g})]} = K_{a,\text{HNO}_3} / \left(1 + \frac{\rho_w}{H_{\text{HNO}_3} \cdot R \cdot T \cdot \text{AWC}}\right) \quad (5-3)$$

where K_{a,NH_3} and K_{a,HNO_3} are the classical dissociation constants of NH_3 and HNO_3 (mol kg^{-1}), ρ_w is the density of liquid water ($1 \times 10^{12} \text{ } \mu\text{g m}^{-3}$), R is the universal gas constant ($8.205 \times 10^{-2} \text{ atm L mol}^{-1} \text{ K}^{-1}$), T is the absolute temperature (K), AWC is the aerosol water content ($\mu\text{g m}^{-3}$), and H_{NH_3} and H_{HNO_3} are the Henry's law coefficients ($\text{mol kg}^{-1} \text{ atm}^{-1}$).



For ambient aerosols, non-ideality due to elevated ionic strength should be considered, as it directly influences the calculation of the effective acid dissociation constant. Accounting for this through the inclusion of activity coefficients, the expression for the effective acid dissociation constant in a non-ideal multiphase system is given as:

$$K_{a,NH_3}^{*,ni} = K_{a,NH_3} \cdot \left(1 + \frac{\rho_w}{H_{NH_3} \cdot R \cdot T \cdot AWC}\right) \cdot \frac{\gamma_{NH_4^+}}{\gamma_{H^+}} \quad (6-1)$$

$$K_{a,HNO_3}^{*,ni} = K_{a,HNO_3} / \left(1 + \frac{\rho_w}{H_{HNO_3} \cdot R \cdot T \cdot AWC}\right) / (\gamma_{NO_3^-} \cdot \gamma_{H^+}) \quad (6-2)$$

Finally, the buffer capacities associated with the NH_4^+/NH_3 and HNO_3/NO_3^- conjugate pairs are expressed as:

$$\beta_{NH_3} = 2.303 \frac{K_{a,NH_3}^{*,ni} \cdot [H^+]}{(K_{a,NH_3}^{*,ni} + [H^+])^2} \cdot [NH_3(g) + NH_3(aq) + NH_4^+(aq)]_{tot}^* \quad (7-1)$$

$$\beta_{HNO_3} = 2.303 \frac{K_{a,HNO_3}^{*,ni} \cdot [H^+]}{(K_{a,HNO_3}^{*,ni} + [H^+])^2} \cdot [NH_3(g) + NH_3(aq) + NH_4^+(aq)]_{tot}^* \quad (7-2)$$

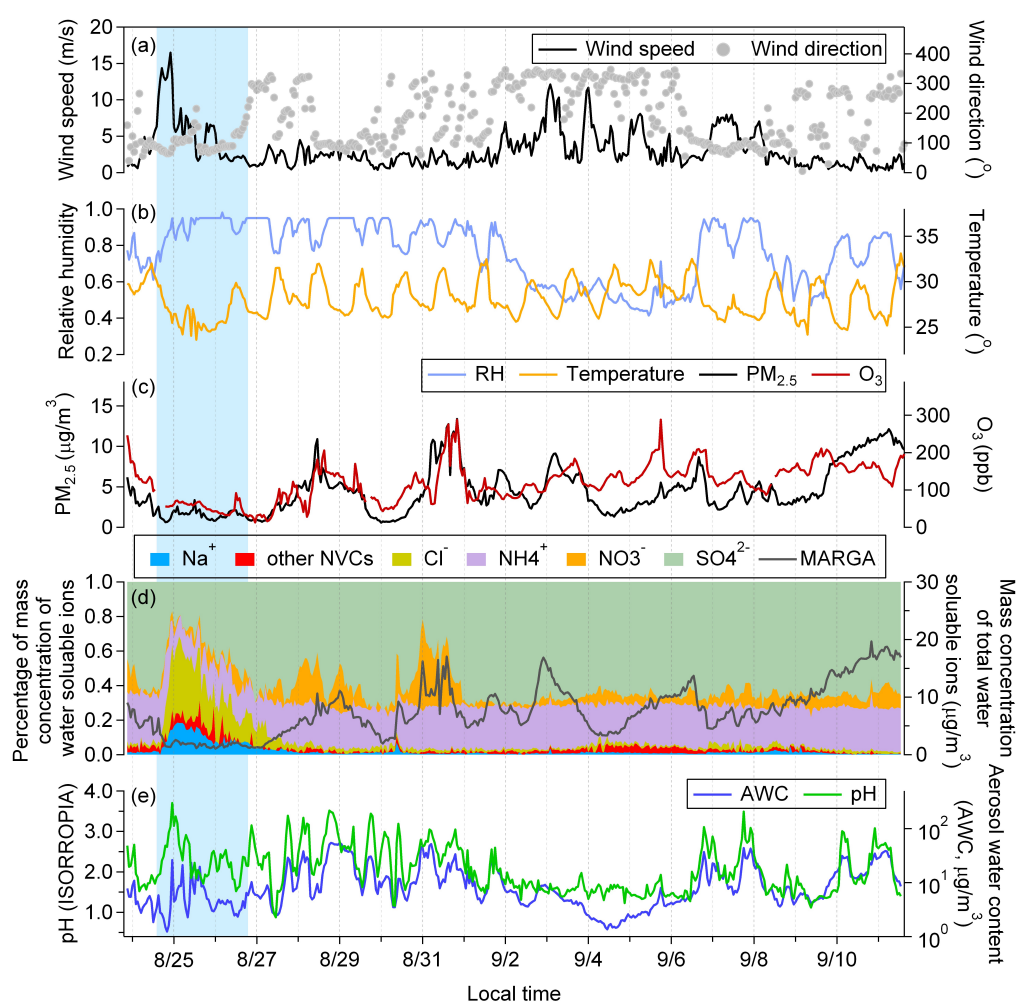
Multiphase buffer theory thus establishes a quantitative framework linking aerosol acidity to multiphase equilibria. Further details of the theoretical derivation are presented in the Supporting Information (SI, Text S1 and Table S1).

3 RESULTS AND DISCUSSION

3.1 Buffer Capacity for Aerosol Multiphase Systems. In this study, MARGA measurements were conducted from 24 August to 10 September 2022 at the Xichong site. Meteorological conditions, chemical compositions obtained from MARGA, and the aerosol water content and pH values simulated by ISORROPIA throughout the sampling period are presented in Figure 1. During the sampling period, Typhoon Ma-on (No. 2209) passed over the sampling site at 15:00 local time on August 24, 2022. Upon its arrival, wind speed at Xichong reached 16.5 m s^{-1} , the highest of the observation period. Concurrent decreases in $PM_{2.5}$, ozone, and total water-soluble ion mass concentrations were observed, accompanied by pronounced increases in the fractional contributions of chloride, sodium, and other NVCs (e.g., Ca^{2+} , Mg^{2+} , K^+) (Figure 1a–d). This



211 interval is hereafter referred to as the typhoon episode (blue shading in Figure 1), with all
212 remaining periods considered non-typhoon episodes. The ISORROPIA-simulated pH values
213 averaged 2.81 ± 0.54 during the typhoon episode and 1.59 ± 0.45 during non-typhoon periods
214 (Figure 1e).



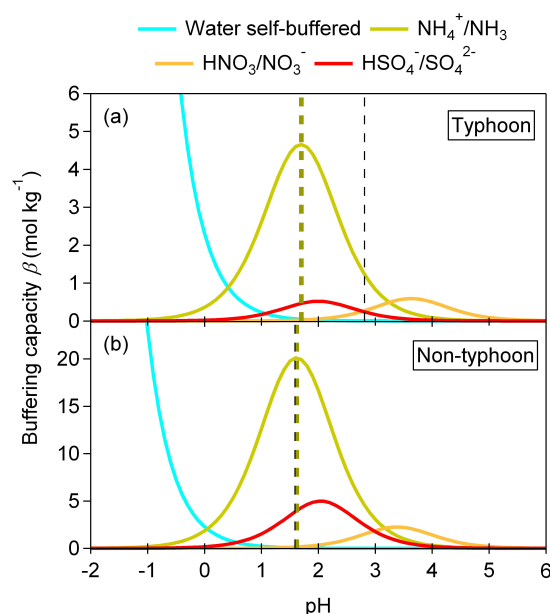
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216 **Figure 1.** Time series of (a) wind direction and wind speed, (b) relative humidity (RH) and
217 temperature, (c) PM_{2.5} and O₃ concentrations, (d) aerosol composition measured by MARGA, and
218 (e) aerosol pH and aerosol water content (AWC) simulated by ISORROPIA at the sampling site.



219 The gray line in panel (c) denotes the total concentration of water-soluble ions measured by
220 MARGA. Other NVCs include Ca^{2+} , Mg^{2+} , and K^{+} . The blue shading indicates the typhoon period.

221 We applied the multiphase buffer theory to calculate the buffering capacities (β) of individual
222 buffering agents ($\text{NH}_4^{+}/\text{NH}_3$, $\text{HNO}_3/\text{NO}_3^{-}$, and $\text{HSO}_4^{-}/\text{SO}_4^{2-}$) under both typhoon and non-typhoon
223 scenarios. In both scenarios, the largest buffering capacity was associated with the $\text{NH}_4^{+}/\text{NH}_3$ pair,
224 followed by $\text{HSO}_4^{-}/\text{SO}_4^{2-}$ and $\text{HNO}_3/\text{NO}_3^{-}$ (Figure 2). The peak buffer pH (defined as the pH
225 corresponding to the highest local maximum of β) for the non-typhoon scenario was ~ 1.62 (Figure
226 2a), closely matching the ISORROPIA-modeled result (1.59 ± 0.45). In contrast, the peak buffer
227 pH under typhoon scenario was ~ 1.70 (Figure 2b), showing a larger discrepancy from the
228 ISORROPIA-simulated result (2.81 ± 0.54).



229

230 **Figure 2.** Buffering capacity (β) of the aerosol multiphase system for typhoon (a) and non-typhoon
231 (b) scenarios. The olive dashed line indicates the pH corresponding to the highest local maximum
232 of β , while the black dashed line represents the ISORROPIA-simulated pH.



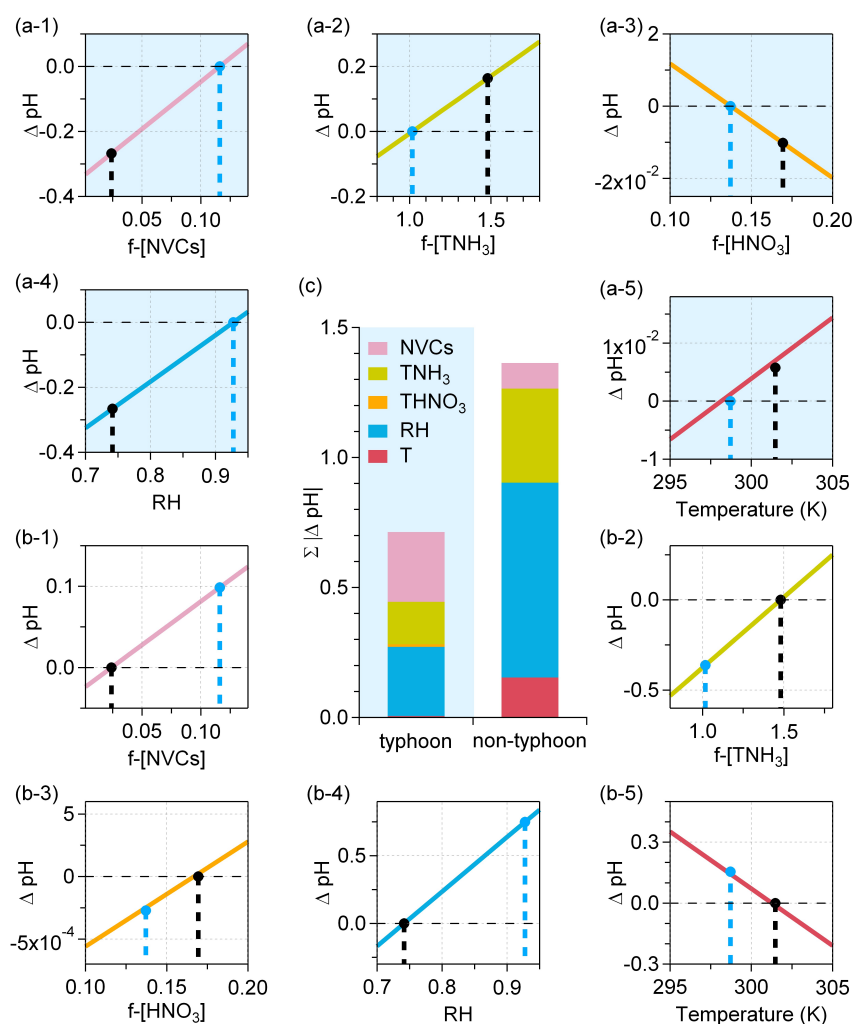
233 The stronger role of $\text{NH}_4^+/\text{NH}_3$ under non-typhoon conditions reflects the abundance of
234 ammonia in this coastal environment and its efficient partitioning, which provides stable buffering
235 at acidic pH. However, for the typhoon scenario, the multiphase buffer theory appears to be less
236 applicable. It should also be noted that multiphase buffer theory assumes instantaneous
237 thermodynamic equilibrium and primarily considers inorganic conjugate acid-base pairs, while the
238 potential contributions of organic species and kinetic effects are neglected. Such simplifications
239 may further contribute to discrepancies under dynamic conditions.

240 **3.2 Contribution of Individual Drivers.** We further quantified the changes in aerosol pH (ΔpH)
241 between typhoon and non-typhoon scenarios attributable to individual drivers (Figure 3), including
242 anion-normalized nonvolatile cations (NVCs) and total ammonia (TNH_3), the fraction of total
243 nitric acid (THNO_3) in anions, relative humidity (RH), and temperature (T). It should be noted that
244 the perturbation constraints imposed on each driver differ both in magnitude and physical meaning
245 (i.e., x-axis criteria in Figure 3, see Text S2 in SI). As a result, the ΔpH contributions of different
246 drivers within the same scenario are not strictly comparable. Figure 3c presents the ΔpH
247 contributions using absolute bar charts, which highlight their absolute magnitudes rather than
248 relative shares. In contrast, the ΔpH contributions of a given driver between the typhoon and non-
249 typhoon scenarios are derived under the same constraint framework and can therefore be
250 meaningfully compared. In Figure 3, the blue and black dashed lines represent the corresponding
251 values for typhoon and non-typhoon scenarios, respectively. For each driver, the ΔpH values
252 correspond to the differences between the two scenarios used as constraints. The constraint ranges
253 were selected to reflect the observed variability during the campaign, thereby ensuring that the
254 perturbations remained within realistic atmospheric conditions.

255 The results indicate that RH, TNH_3 , and temperature contributed substantially more to ΔpH in
256 the non-typhoon scenario than in the typhoon scenario, with 0.75 vs. 0.26 units for RH, 0.36 vs.
257 0.16 units for TNH_3 , and 0.15 vs. 0.01 units for temperature, respectively. Notably, NVCs was the
258 only driver exhibiting a larger contribution to ΔpH in the typhoon scenario than in the non-typhoon
259 scenario (0.27 vs. 0.09 units). In other words, the contribution of chemical components,



260 particularly NVCs, to ΔpH was more pronounced in the typhoon scenario, which partly explains
261 the discrepancy between the pH predicted by the multiphase buffer theory and that calculated by
262 ISORROPIA. This shift highlights a transition from meteorologically driven controls (RH and
263 temperature) under non-typhoon conditions to compositionally driven controls dominated by sea-
264 salt derived cations during the typhoon.



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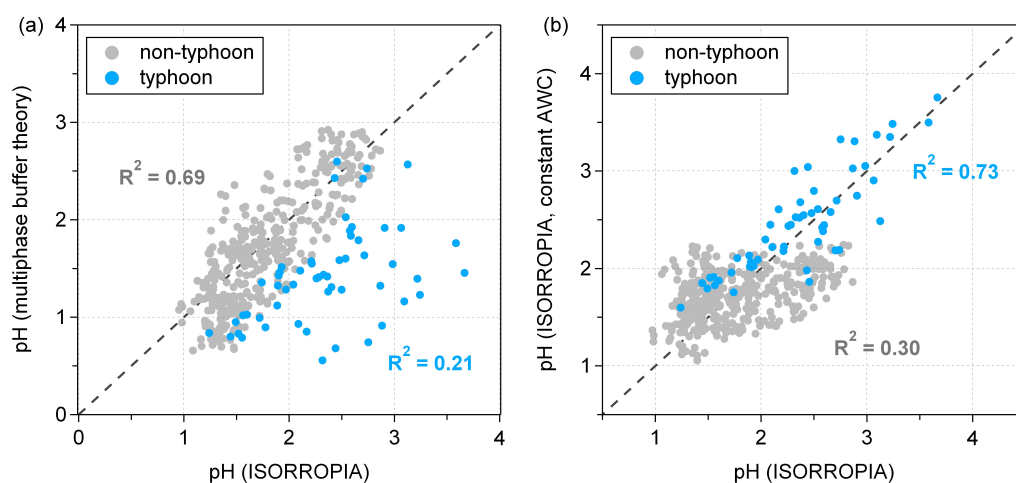
Figure 3. Aerosol pH differences (ΔpH) between typhoon (a1–5) and non-typhoon (b1–5) scenarios, attributed to anion-normalized non-volatile cations (NVCs) and total $[\text{NH}_3]$ (TNH_3), the fraction of total $[\text{HNO}_3]$ (THNO_3) in anions, relative humidity (RH), and temperature (T). The light blue background indicates simulations using typhoon conditions as inputs (a1–5), whereas the white background denotes simulations with non-typhoon conditions (b1–5). Panel (c) summarizes the total absolute ΔpH ($\Sigma|\Delta\text{pH}|$) of different drivers in both scenarios. The blue and black dashed lines represent the corresponding values for typhoon and non-typhoon scenarios, respectively

Previous studies in inland regions (Zheng et al., 2020), such as the North China Plain and the southeastern United States, have emphasized ammonia availability and aerosol water as the primary determinants of aerosol acidity, whereas our results underscore the unique role of NVCs in coastal megacities under extreme weather influence. For $\text{NH}_4^+/\text{NH}_3$ to serve as the dominant buffering pair, two conditions must be satisfied (Zheng et al., 2020): (1) the equivalent charge of total cations exceeds that of total anions, and (2) the equivalent charge of NVCs is smaller than that of nonvolatile anions. In the typhoon scenario, however, the equivalent charge of NVCs exceeded that of nonvolatile anions, thereby violating the conditions required for $\text{NH}_4^+/\text{NH}_3$ to serve as the dominant buffering pair. This explains why the multiphase buffer theory fails to reliably predict aerosol pH under typhoon conditions. The enhanced influence of NVCs can be attributed to the substantial influx of sea-salt particles transported by strong winds and the altered trajectories of air masses during the typhoon. These inputs directly neutralize acidic species and disrupt the conventional $\text{NH}_4^+/\text{NH}_3$ buffering system, a process far less pronounced in inland settings.

We compared the ISORROPIA-simulated pH with that predicted by the multiphase buffer theory and with the ISORROPIA-simulated pH under a constant-AWC assumption, for typhoon and non-typhoon scenarios, respectively (Figure 4). The constant-AWC experiment was designed to isolate the role of chemical composition from that of aerosol water content, thereby allowing us to directly assess whether pH variability can be reproduced without explicitly accounting for



dynamic changes in AWC. The results show that under the non-typhoon scenario, ISORROPIA-simulated pH gives a stronger correlation with the multiphase buffer theory ($R^2 = 0.69$) than with the constant-AWC simulation ($R^2 = 0.30$). In contrast, under the typhoon scenario, ISORROPIA-simulated pH correlates well with constant-AWC ($R^2 = 0.73$), but poorly with the multiphase buffer theory ($R^2 = 0.21$). These findings indicate that in regimes or scenarios buffered by the $\text{NH}_4^+/\text{NH}_3$ pair, such as non-typhoon conditions in Shenzhen, variations in AWC alone can provide a reliable prediction of aerosol pH, even without explicitly accounting for the temporal and spatial variability in particle chemical composition. The buffering effect of ammonia suppresses the influence of compositional differences, making aerosol water content the primary determinant of aerosol pH. However, in environments that are not buffered by $\text{NH}_4^+/\text{NH}_3$, the influence of AWC on pH becomes weaker, and reliable predictions require explicit consideration of chemical composition within a constant-AWC framework.



305

Figure 4. Correlations between aerosol pH simulated based on ISORROPIA and pH predicted by (a) the multiphase buffer theory, and (b) ISORROPIA under a constant aerosol water content (AWC) assumption with varying chemical compositions. Blue and gray dots represent typhoon and non-typhoon scenarios, respectively. The black dashed line indicates the 1:1 reference.



310 These results suggest that chemical composition plays a more critical role in determining pH
311 under typhoon scenarios, whereas during non-typhoon periods, aerosol water content influenced
312 by RH and temperature exerts a stronger impact on pH. While the multiphase buffer theory is
313 robust under ammonia-buffered regimes, its applicability becomes limited in environments
314 dominated by NVCs. Refining the framework to explicitly incorporate composition-meteorology
315 interactions is therefore essential for accurately predicting aerosol acidity in coastal regions subject
316 to extreme weather events.

317 **3.3 Atmospheric implications.** In densely populated continental regions, where anthropogenic
318 emissions and atmospheric ammonia concentrations are high, aerosol pH is likely controlled by
319 the $\text{NH}_4^+/\text{NH}_3$ buffering pair and can therefore be reasonably approximated on the basis of aerosol
320 mass concentration and liquid water content. This provides an opportunity to reconstruct long-term
321 trends and large-scale spatial distributions of aerosol pH, and implies that emission reductions
322 targeting ammonia and sulfate can exert direct and predictable influences on aerosol acidity.

323 In coastal regions, however, whether ammonia serves as the dominant buffering pair is strongly
324 influenced by seasonality and meteorological conditions. Our field observations directly captured
325 two contrasting situations: a non-typhoon period, in which ammonia acted as the dominant
326 buffering pair, and a typhoon period, during which NVCs played a more important role. In the
327 former case, aerosol pH can be reliably predicted using only aerosol mass concentration and AWC,
328 whereas in the latter case, more detailed chemical composition information is required for accurate
329 prediction. Acidity regulation in coastal area is inherently more variable than in continental settings,
330 with rapid transitions between meteorological and compositional control. Such variability
331 complicates the prediction of secondary aerosol formation and the assessment of pH-dependent
332 processes (e.g., metal solubility and heterogeneous chemistry). Our results therefore underscore
333 the need for improved representation of composition-meteorology interactions in chemical
334 transport models and highlight the need for targeted observations during extreme weather events
335 to constrain acidity regulation in coastal atmospheres.

336



337 **Data availability.** Data used to produce the plots within this work are available in Zenodo ([https://](https://zenodo.org/records/17207845)
338 <https://zenodo.org/records/17207845>).

339 **Author contributions.** JZ and XY designed the study. JZ and YZ analyzed the data. JZ wrote the
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341 **Competing interests.** The contact author has declared that none of the authors has any competing
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