



Coupled C and N turnover in a dynamic pore scale model reveal the impact of exudate quality on microbial necromass formation

Maximilian Rötzer^{1,4}, Henri Braunmiller², Eva Lehdorff³, Nadja Ray^{1,4}, Andrea Scheibe³, and Alexander Prechtel¹

¹Modeling and Numerics, Department of Mathematics, University of Erlangen–Nürnberg, Erlangen, Germany

²Root-Soil Interaction, School of Life Sciences, Technical University of Munich, Freising, Germany

³Soil Ecology, Bayreuth Center of Ecology and Environmental Research (BayCEER), University of Bayreuth, Bayreuth, Germany

⁴Mathematical Institute for Machine Learning and Data Science, Catholic University of Eichstätt-Ingolstadt, Ingolstadt, Germany

Correspondence: Maximilian Rötzer (maximilian.roetzer@fau.de)

Abstract. The adequate quantification of soil organic carbon (SOC) turnover is a pressing need for improving soil health and understanding climate dynamics. It is controlled by the complex interplay of microbial activity, availability of carbon (C) and nitrogen (N) sources, and the dynamic restructuring of the soil's architecture. Accurate modeling of SOC dynamics requires the representation of these processes at small spatial scales.

- 5 We present a mechanistic, spatially explicit model at the pore scale, which couples enzymatic degradation of particulate organic matter (POM), microbial necromass and root exudates with microbial growth and turnover, C respiration and N cycling depending on the C/N ratios of the different organic carbon sources. It is combined with a cellular automaton model for simulating soil structure dynamics including the stabilization of soil particles, POM or microbial necromass via organo-mineral associations.
- 10 The virtual soil simulations use μ CT data of aggregates and parameters from rhizosphere experiments without parameter fitting to explore the influence of (i) soil structural heterogeneity and connectivity, (ii) N limitation, and (iii) necromass formation on SOC storage.

- Our results highlight that evolving soil architecture and pore connectivity control substrate accessibility, creating micro-scale hot and cold spots for microbes. N availability consistently co-limits microbial growth, while a favorable C/N ratio of root exudates substantially reduces respiration and increases CUE over extended periods. Necromass emerges as long-term SOC pool, as N from short-term root exudation pulses promotes biomass growth and is converted into slowly degradable necromass, which can be physically protected through occlusion. The findings align with lab experiments and additionally allow us to elucidate the spatial and temporal dynamics of the drivers of carbon turnover.
- 15

1 Introduction

- 20 Soil organic carbon (SOC) turnover is a fundamental component of the global carbon cycle, influencing soil health and climate dynamics. In their review Vogel et al. (2024) emphasized the importance of explicit representation of biological processes in



models to adequately describe soil functions such as the turnover of organic matter, N cycling, or soil structure formation, and state that this is a prerequisite to arrive at the urgently needed predictive models understanding perturbations caused by soil management or current changes in climatic conditions. Varney et al. (2022) compared soil carbon simulations in current earth system models and concluded that much of the uncertainty can be attributed to the simulation of below-ground processes, and greater emphasis is required on improving their representation in future developments of models. Accurate modeling of SOC dynamics requires a nuanced understanding of a range of interconnected biological and physical processes occurring at small spatial scales (He et al., 2024; Pot et al., 2021).

We aim to address several challenges: First, the highly heterogeneous soil structure and dynamically evolving connectivity intricately weave together microbial activities and stabilization mechanisms. On the small scale, aggregates can be seen as microbial reactors whose evolving properties should be accounted for in bottom-up modeling (Dungait et al., 2012; Wang et al., 2019). The microbial turnover of organic matter (OM), the subsequent formation of gluing spots and the build-up of organo-mineral associations are major aspects in soil structure formation, and may on the other hand lead to physical protection of particulate organic matter (POM) (Totsche et al., 2018). Thus, structural dynamics and biological processes cannot be considered independently (Philippot et al., 2024).

Second, the soil-microbe system responds to different slow or fast-cycling carbon/nitrogen sources. Besides plant-derived POM, microbial necromass can contribute to a significant proportion of SOC and must be explicitly taken into account (Chandel et al. (2023); Liang et al. (2019); Kästner et al. (2021)). In the rhizosphere root exudates represent an important and easily accessible carbon source for microbes (Baumert et al., 2018). Mainly due to the limitation of experimental techniques, the quantification of their impact on aggregation and carbon turnover in soils remains difficult (Gregory et al., 2022).

Third, the availability of resources and the specific microbial C/N demand determine microbial growth, whereas the microbial decomposition again feeds back on the overall C and N distribution in soils (Zechmeister-Boltenstern et al., 2015; Simon et al., 2024). This interplay should be included in process-based models to better understand the dynamic nutrient cycling and microbial C/N efficiency in soils (Kaiser et al., 2014; Zechmeister-Boltenstern et al., 2015; Sokol et al., 2022). He et al. (2024) emphasized that the link between microbial carbon use efficiency (CUE) and SOC is mediated by processes that stabilize microbial necromass within soil aggregates or associate it with mineral surfaces.

The complexity of soil systems needs mathematical approaches to study the interplay of the underlying mechanisms and make more precise predictions of CUE responses under various environmental scenarios. As pointed out in several reviews (Baveye et al., 2018; Pot et al., 2021; Vogel et al., 2022), only explicit spatial and image-based modeling at the pore scale, and not pool or lumped parameter models are capable to examine the interaction between architecture, soil microbial activity and OM turnover.

The variability in soil structure as a consequence of the self organization of soil aggregates was first modeled by Crawford et al. (1997) and pursued further simulating aggregate formation as a function of binding surfaces (of chemical or microbial origin) in Crawford et al. (2012). Ray et al. (2017) and Rupp et al. (2018b) developed a comprehensive cellular automaton method (CAM) to simulate aggregate formation in a spatially and temporally explicit way. Zech et al. (2022b) applied a CAM



to quantify the simultaneous structure formation/disaggregation and OM turnover and study the drivers texture, decomposition rate, and POM input. They revealed that the addition of fresh organic matter to soil may induce a short-term primed decomposition of native OM due to the creation of new gluing spots leading to the simultaneous break-up of older organo-mineral associations and the subsequent decomposition of POM.

60 The first modeling study to combine root growth and exudation, microbial growth, particle surface alterations, aggregate dynamics and C turnover on the pore scale within a CAM framework was Rötzer et al. (2023). Their results support that exudates, although degraded within a few days, are important drivers for structure formation, leading to a more persistent stabilization of aggregate structures via an increased amount of gluing spots even after root degradation than in comparison to aggregate structures developing in the presence of only POM.

65 Portell et al. (2018) presented an individual-based model using CT scans of natural soils to study the effects of heterogeneously distributed organic matter on different strains of bacteria at the mm scale.

Zech et al. (2020) investigated the role of connectivity for local and global CO₂ output due to the hydrolysis of POM and bacteria dynamics in unsaturated, heterogeneous micro-aggregates.

In the present work we extend the pore scale model for coupled carbon turnover, microbial dynamics and structure evolution
70 (Zech et al., 2022b; Rötzer et al., 2023) to account for different carbon sources and include the interaction between C and N cycles. In the experimental limitation of spatially identifying the heterogeneity of C/N ratios in microbially active zones in soils, such as the rhizosphere and equivalent inactive zones, this simulation approach makes use of the current knowledge about coupled C and N cycling to point out where OM stabilization processes in the form of microbial necromass potentially occur and provide first quantitative data. Therefore, we employ a biogeochemical model to define the demand for C and N of
75 the microbial biomass as delineated in (Manzoni and Porporato, 2009; Manzoni et al., 2012; Kaiser et al., 2014):

1. POM and microbial necromass are enzymatically degraded while forming reactive organo-mineral associations and then contribute to dissolved C and N in the soil solution. The enzymatic degradation is enabled by microbial biomass.
2. Microbial biomass is degraded to microbial necromass while the C/N ratio of 10 is constant. Moreover, it enhances the formation of gluing spots on mineral surfaces.
- 80 3. Root exudates directly contribute to DOC and DON. High concentrations of such remnants in the soil solution likewise induce the creation of gluing spots temporarily coating mineral surfaces.
4. Gluing spots facilitate the aggregation of primary particles, microbial necromass, POM and soil micro-aggregates.
5. DOC and DON are taken up by microbial biomass and assimilated in the optimal C/N ratio of 10 for growth and maintenance. During all these processes C is respired as CO₂ defining the variable CUE.
- 85 6. As the uptake of C and N of the microbial biomass may not be balanced, the surplus of C is leaving the system. A surplus of N gets recycled as dissolved N. Furthermore a portion of dissolved N is leached over time.

The interdependence of the different model components and the transformation between them are depicted in Fig. 1.

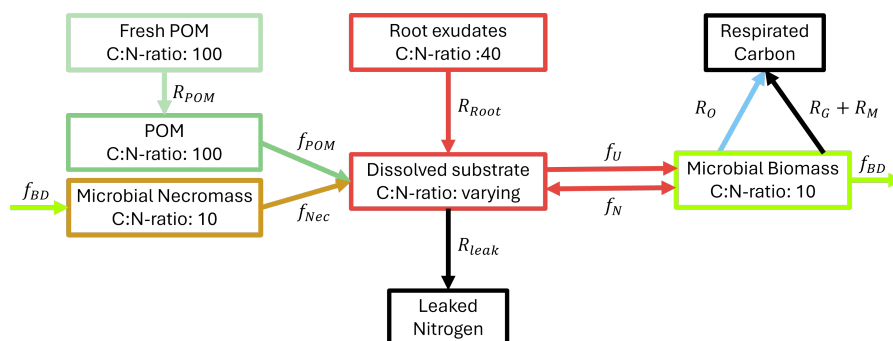


Figure 1. Conceptual model of the microbial dynamics with distinct carbon and nitrogen pools and fluxes between them.

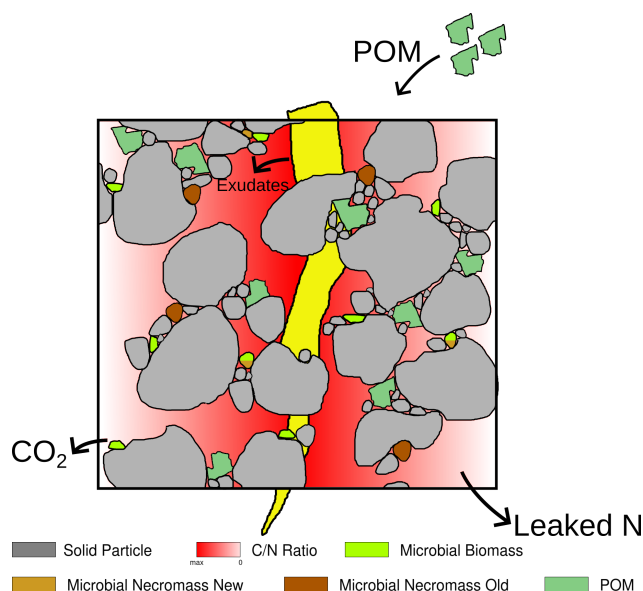


Figure 2. Illustration of the components of the model. The drawing includes a root (yellow) which is not taken into account explicitly as a model component, only its exudates.

The microbial turnover of slowly decaying POM, microbial necromass and easily available carbon sources as exudates with different C/N ratios and concentrations will be studied in varying scenarios. A hypothetical root in the center of the domain exudes and contributes to the dissolved C and N concentrations, cp. the conceptual drawing in Fig. 2. Organo-mineral interactions are taken into account including the dynamic (re)aggregation of soil particles so that SOC protection or priming effects are captured by the explicit resolution of the pore scale dynamics. The CUE is a dynamic quantity resulting from the environmental conditions and can be tracked and studied over the simulations.

The aim is to study the influence of N dynamics on microbial C turnover and to evaluate the importance of necromass as SOC pool, both interacting with soil structure on the pore scale.



2 Materials and methods

We introduce a process-based, spatially and temporally explicit micro-scale model that describes C and N fluxes in soils along with the decomposition of particulate organic matter (POM), the development of microbial bio- and necromass, and simultaneous aggregate turnover. The latter is a consequence of the surface properties of soil particles, which are altered by the production of gluing spots by microbes and organic remnants. The dynamic re-arrangement of primary soil particles and soil micro-aggregates impacts the spatial accessibility of OM and consequently its degradation rates. We also account for the input of root exudates as an easily available carbon source.

2.1 Cellular automaton and soil restructuring

To account for the spatial distribution and re-arrangement of model components, we conduct our investigations in the flexible framework of a cellular automaton method extending the work of Ray et al. (2017); Rupp et al. (2019); Zech et al. (2022b) and Rötzer et al. (2023). Within the CAM, the computational domain is discretized into voxels (cells) of $2\mu\text{m} \times 2\mu\text{m} \times 1\mu\text{m}$, the smallest entities of the model. They have different (possibly co-existing) states such as solid, pore, POM, biomass or necromass, see Fig. 3. The solid phase consists of water-stable, inseparable, but mobile primary soil particles and micro-aggregates, see Appendix Section A1 for a detailed description. The pore space is assumed to be either saturated with soil solution carrying the dissolved substrates carbon and nitrogen, or occupied by solid microbial necromass or POM. They may also co-exist. A key component of the model is the (dis-)aggregation of mobile particles according to different surface attractions. This concept is elaborated in detail in Rötzer et al. (2023) and Zech et al. (2022b), see also the Appendix. This attractiveness includes that with increasing size, a decreasing fraction of the soil particles' surfaces is marked as reactive. At these reactive surfaces, POM particles may attach preferably undergoing microbial degradation (Witzgall et al., 2021). The concept of reactive surfaces thus includes gluing spots originating from microbial remnants and occurring on temporary OM covered mineral surfaces which enhance the formation of aggregates (Bucka et al., 2019). As this surface conditioning of mineral surfaces by organic compounds may be locally retained even after degradation of the OM coating (Kleber et al., 2007), the term memory edges for mineral surfaces after degradation of OM has been introduced in Zech et al. (2022b). By interacting with the mineral surfaces these gluing spots strengthen the adhesive bonds between soil particles. Following Nazari et al. (2022), extracellular polymeric substances (EPS) and root exudates, such as mucilage, follow the same principle in enhancing the adhesion of mineral surfaces in soil. As they are both gelatinous high-molecular-weight substances, both substances create equally favorable hot-spots for the development of gluing surfaces.

2.2 C and N turnover

We extend the models in Zech et al. (2022b) and Rötzer et al. (2023) to incorporate N fluxes, take into account microbial growth dependent on C and N availability, and include necromass evolution as an important OM pool.

Following the review by Manzoni and Porporato (2009) and Kaiser et al. (2014) on C/N pathways in soils, we consider the following model components directly related to different C and N pools with specific C/N ratios:

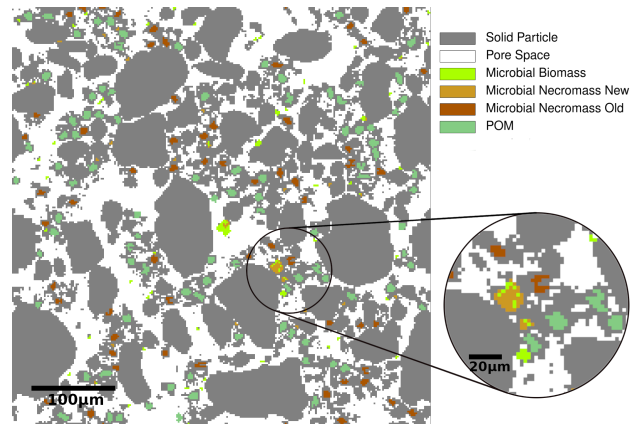


Figure 3. Sketch of the 2D discretized computational domain of 250x250 pixels with a resolution of 2 μm and different cell types. Edge types are not displayed for the ease of presentation.

130

- Particulate organic matter (POM), rich in C with $C/N=100$
- Microbial biomass with $C/N=10$
- Microbial necromass with $C/N=10$
- Root exudates with $C/N=40$
- Soil solution with dissolved C and N (dynamic C/N ratio)
- C and N sinks due to respiration and leakage

135

We follow a concentration-based approach and, opposed to Manzoni and Porporato (2009), combine the N assimilation pathways, the direct organic uptake and mineralization–immobilization.

We assume that the pools microbial biomass, necromass, and POM are strictly homeostatic, i.e. they keep their C/N ratios constant over time. Contrarily, dissolved C and N can be assimilated and diffuse independently from each other, i.e. the C/N ratio there is variable.

140

We now state the underlying model equations - a system of differential equations - describing the evolution of the different C and N concentrations in the respective C and N pools and the fluxes between these pools as depicted in Fig. 1. The unknowns and reaction rates are summarized in Table 1 and 2 and the related parameters are specified in Table 4.



Table 1. State Variables (unknowns) of the model.

Variable	Description
C_B	C in microbial biomass
C_{Nec}	C in microbial necromass
C_{CO_2}	Accumulated C from CO_2 respired by the microbes
C_{POM}	C in particulate organic matter
C_S	Organic C substrate dissolved in soil solution (DOC)
N_B	N in microbial biomass
N_{leak}	Accumulated leaked N, no re-entry into the system
N_{Nec}	N in microbial necromass
N_{POM}	N in particulate organic matter
N_S	N dissolved in soil solution

Table 2. Reaction rates of the model and their definition. The fluxes f quantify the transformation of C and N from one pool to another while the rates denoted by R represent a (external) loss or source, accounting for processes such as respiration, leakage, fresh POM input or root exudation.

Rate	Description
$f_U = \nu_{CS} \frac{C_S}{C_S + k_{CS}} C_B$	Microbial uptake of C
$f_N = \frac{f_U - R}{(C/N)_B}$	Microbial assimilation or mineralization of N
$f_{POM} = k_{POM} \left(0.5 + \frac{C_B^{total}}{C_B^{init}} \right) \frac{s_f^p}{s^p} C_{POM}$	Decomposition of POM
$f_{Nec} = k_{Nec} \left(0.5 + \frac{C_B^{total}}{C_B^{init}} \right) \frac{s_f^p}{s^p} C_{Nec}$	Decomposition of necromass
$f_{BD} = k_{BD} C_B$	Microbial decay
$R_O = \partial_t (C_B - (\frac{C}{N})_B N_S)$ or 0	Overflow respiration, see Appendix Eq. (A2)
$R_G = k_G f_U$	Growth respiration
$R_M = k_M C_B$	Maintenance respiration
$R = R_G + R_M + R_O$	Total respiration
$R_{Leak} = k_{leak} N_S$	Leaked N
R_{POM}	Discrete input in time: addition of fresh POM particles
R_{Root}	Input function of time for root exudation



$$\frac{\partial C_B}{\partial t} = f_U - R_G - R_M - f_{BD} - R_O \quad (1)$$

$$\frac{\partial N_B}{\partial t} = \frac{\partial C_B}{\partial t} \cdot \frac{1}{(C/N)_B} \quad (2)$$

$$145 \quad \frac{\partial C_S}{\partial t} - \nabla \cdot (D_S \nabla C_S) = f_{POM} + f_{Nec} - f_U + R_{Root} \quad (3)$$

$$\frac{\partial N_S}{\partial t} - \nabla \cdot (D_S \nabla N_S) = \frac{f_{POM}}{(C/N)_{POM}} + \frac{f_{Nec}}{(C/N)_{Nec}} - f_N + \frac{R_{root}}{(C/N)_{root}} - R_{leak} \quad (4)$$

$$\frac{\partial C_{Nec}}{\partial t} = f_{BD} - f_{Nec} \quad (5)$$

$$\frac{\partial N_{Nec}}{\partial t} = \frac{f_{BD}}{(C/N)_B} - \frac{f_{Nec}}{(C/N)_{Nec}} \quad (6)$$

$$\frac{\partial C_{POM}}{\partial t} = -f_{POM} + R_{POM} \quad (7)$$

$$150 \quad \frac{\partial N_{POM}}{\partial t} = \frac{-f_{POM}}{(C/N)_{POM}} + \frac{R_{POM}}{(C/N)_{POM}} \quad (8)$$

$$\frac{\partial C_{CO_2}}{\partial t} = R_G + R_M + R_O \quad (9)$$

$$\frac{\partial N_{leak}}{\partial t} = R_{leak} \quad (10)$$

The reaction and transport system Eq. (1)– Eq. (10) determines how C and N cycles between microbial biomass (C_B and N_B), microbial necromass (C_{Nec} and N_{Nec}), dissolved substrates in pore space (C_S and N_S) and organic soil compounds
155 such as particulate organic matter (C_{POM} and N_{POM}). We also track the amount of C that is transformed to CO_2 as C_{CO_2} and the leaked N as N_{leak} .

Following the review of Manzoni and Porporato (2009), in which a theoretical framework of well-established microbial biomass balance equations for the soil substrate–decomposer dynamics considering both, the microbial stoichiometry and the impact of limiting nutrients, is introduced, the evolution of C in biomass, Eq. (1), depends on the C uptake from soil solution
160 (f_U), the growth (R_G) and maintenance (R_M) respiration, the decay to necromass (f_{BD}), and a potential overflow mechanism (R_O). The flux f_U describes the amount of C taken up from soil solution by the microbial biomass, modeled via a Michaelis-Menten kinetics. The growth respiration rate R_G is assumed to be proportional to f_U . The basal metabolic turnover of microbes is summarized as maintenance respiration R_M and is modeled as first-order kinetics. The decay term f_{BD} accounts for the limited lifetime of microbial biomass and its transformation into microbial necromass thereafter. The overflow rate R_O finally
165 reflects the release of excess C from the system in the case of N limitation, see the Appendix, Section A4 for a more detailed description.

Diffusion–reaction equations for C_S , Eq. (3), and N_S , Eq. (4), describe the uptake from, release into and diffusive transport of C and N inside the soil solution. We account for sources, specifically root exudation R_{Root} , hydrolysis of POM f_{POM} and necromass f_{Nec} and sinks via microbial uptake f_U and leaked dissolved N R_{leak} .



Table 3. Particle size distribution of the SPE loam.

Feret diameter	1-2 μm	2-6 μm	6 -20 μm	20 - 63 μm	63 - 200 μm
Contribution	19.5 %	5.2 %	16.4 %	26.3 %	32.6 %

170 In Eq. (4), the term f_N can be positive or negative, representing the assimilation or mineralization of N_S as necessary to balance the C demand of the microbial biomass respecting its constant C/N ratio, see discussion above.

In detail, assuming that exoenzymes released by the microbes are ubiquitous in soils the enzymatic hydrolysis of POM f_{POM} and necromass f_{Nec} follow first-order kinetics. Here, the turnover coefficients were modified to represent two decomposition processes suggested by (Chandel et al., 2023): Firstly, the rate coefficients are a function of microbial biomass. Assuming a
 175 basal turnover rate k_{POM} and k_{Nec} , which include physical and chemical processes, the decay rate depends linearly on the total amount of microbial biomass in the domain C_B^{total} , normed by a factor C_B^{init} , the initial total amount of microbial biomass. Secondly, we take into account that the degradation requires access of the POM particles to the pore space, and that the lack of this access scales down the rate (physical occlusion, Zech et al. (2022b)). In terms of the CAM concept (see Section A1), we evaluate the ratio of the number of surface edges of a POM particle next to a pore cell s_f^p and the total number of edges of the
 180 POM particle s^p . The more a POM particle is occluded, the slower its degradation. A complete occlusion results in the total halt of decomposition, as $s_f^p = 0$. The degradation of necromass is treated identically.

R_{root} describes an input rate of dissolved substrates from root exudation. Finally, a first-order kinetics describes the irreversible loss of dissolved N through leaching R_{leak} .

2.3 Virtual Soil and Model Parametrization

185 The parametrization of the model is oriented at the experimental platforms presented in (Vetterlein et al., 2021), where a multitude of studies has been done at the same sites or with the same soils. We create a virtual soil representing the loam of the soil plot experiment (SPE) which has a bulk density of 1.39 g cm^{-3} and a porosity of 0.411 (Vetterlein et al., 2021), which results in a solid density of 2.36 g cm^{-3} . The particle size distribution of the SPE loam is given in Table 3. Soil particles of realistic shapes are first randomly distributed within a domain of size $(500 \times 500 \times 1) \mu\text{m}^3$ according to the given particle size
 190 distribution and porosity. The shapes of the individual particles stem from a library obtained by dynamic image analysis of wet-sieved soil fractions of an arable Cambiosol with a range of clay contents ((Schweizer et al., 2019); Figure 4a).

The amounts of C and N in pools and mass fluxes are often measured relative to the dry soil mass. The total dry soil mass of our virtual soil can be calculated by the solid density and porosity in the domain, resulting in $0.35 \mu\text{g}$ soil, experimental data can directly be converted and used for our parametrization.

195 As the SPE loam has been used in studies on microbial growth kinetics for different root-derived substrates (Niedeggen et al., 2024) and the quantification of maize root exudation (Santangeli et al., 2024), we gain further relevant parameters as follows (see Table 4 for all values and references).

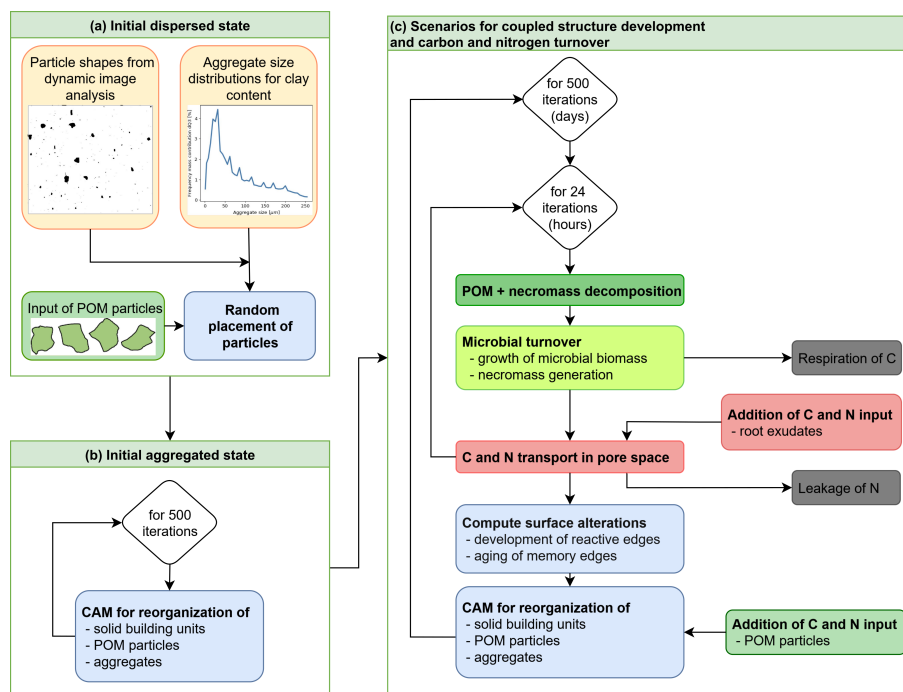


Figure 4. Workflow for the simulations: (a) Creation of initial dispersed state, (b) application of CAM to create initial aggregated state, and (c) carbon and nitrogen turnover coupled with structure development in 500-day scenarios.

Throughout the simulations the pore space is fully saturated with an initial DOC concentration of $C_S^0 = 10^{-5} \text{ gCcm}^{-3}$ and an initial C/N ratio of 10. C enters the domain either as POM particles, or as dissolved C from exudates (see below). Initially
 200 POM particles with a concentration of 0.48 gCcm^{-3} and a C/N ratio of 100, and necromass with a concentration of 0.32 gCcm^{-3} and C/N ratio of 10 amounting to a volume fraction of 5% of the solid area were randomly added to the pore space corresponding to 60 POM and 60 necromass particles each with a size of 6-10 μm in diameter. This results in an initial amount of $6.24 \text{ mgCg}^{-1}\text{POM-C}$ and $2.61 \text{ mgCg}^{-1}\text{microbial-necromass-C}$. While up to 70 % of the initial soil C is retained in POM, its contribution to the initial soil N content is 20 %.

205 After this initialization 216 microbial biomass voxels with a density of 63.4 mgCcm^{-3} , divided into 72 connected components of 3 cells each are distributed in the pore space. This results in a total amount of microbial biomass of $158 \text{ }\mu\text{gCg}^{-1}$, which corresponds to Niedeggen et al. (2024).

In all scenarios we consider a POM input of 0.083 mgCg^{-1} by a random placement of one POM particle with $60 \text{ }\mu\text{m}^2$ into the pore space of our computational domain every 10th day (compare low input scenario in Zech et al. (2022b)). The initial
 210 amount of total organic carbon (TOC) in the system constitutes 0.89 % of the soil's mass, which is close to 0.85 % of Vetterlein et al. (2021), and in the range of other loamy soils (cf., e.g., 0.74% (Endress et al., 2024)). The amount of total organic N (TON) is 0.034 % of soils' mass which could which could correspond to values for crop and grassland soil with low N content, 0.083 % in Vetterlein et al. (2021).



We assume that root exudation takes place along a central vertical line of the computational domain (recall the scheme in Fig. 2). For the root C exudation rate we use the values for maize (wild type) in the plant development stages BBCH14 and 19 reported in Santangeli et al. (2024) and for the Michaelis-Menten kinetics of early season exudates we apply the values reported by Niedeggen et al. (2024), see Table 4.

Further, Niedeggen et al. (2024) reported an initial microbial biomass C_B^0 of $158 \mu\text{gCg}^{-1}$, a microbial basal activity R_{basal} of $0.18 \mu\text{gCO}_2\text{-Cg}^{-1}\text{h}^{-1}$ and a peak activity R_{peak} of $2.54 \mu\text{gCO}_2\text{-Cg}^{-1}\text{h}^{-1}$. Assuming that basal activity is dominated by metabolic activities not related to growth, the maintenance respiration rate can be derived as the fraction between respiration and microbial biomass $k_M = \frac{R_{\text{basal}}}{C_B^0} = 0.3 \cdot 10^{-6} \text{s}^{-1}$.

Similarly assuming that peak respiration is dominated by growth respiration, which depends linearly on the C uptake, the maximal C uptake rate coefficient of a microbial biomass can be derived as $\nu_{\text{CS}} = \frac{(R_{\text{peak}} - R_{\text{basal}})/k_G}{C_B^0} = 1.6 \cdot 10^{-5} \text{s}^{-1}$.

Exudation causes a C release with a rate of $3.6 \frac{\mu\text{gC}}{\text{h}\cdot\text{cm}^2}$ (Santangeli et al., 2024) in the vertical center line (if the cell is not of type soil particle). One pulse of exudation through the surface of $4 \mu\text{m}^2$ results in a total rise of DOC of $0.01 \text{mgCgsoil}^{-1}$ per day, which is the duration of one pulse in the reference setting.



Based on this model parametrization, the simulations are conducted with the algorithmic workflow illustrated in Fig. 4: Starting from the creation of the initial dispersed state with an initial, random placement of the particles, the subsequent creation of the initial aggregated state by the CAM rules, and then the time loop of the different scenarios including the temporal evolution of the components. Time step size for the CAM loop is 1 d, for the calculation of reaction and transport terms 1 h. Details on the CAM and the numerical methods used in the implementation are given in the Appendix A1 and A3.

2.4 Simulation Scenarios

In order to study the C and N dynamics we set up the reference Scenario 1Full (see Section 2.4.1) which is then restricted to Scenario 1Stat by excluding structure dynamics, and extended by varying the quality (Scenario 2CN) and the amount (Scenario 3Pulse) of root exudates to study the influence of different drivers in the system. All simulations cover 500 days.

2.4.1 Scenario 1: C/N dynamics under continuous POM input and a single short-term root exudation

Scenario 1Full: Reference setting

We consider our complete model outlined in Section 2 and parametrized according to Table 4. Within the total simulation time of 500 days, one pulse of root exudates for one day at a rate of $3.6 \frac{\mu\text{gC}}{\text{h}\cdot\text{cm}^2}$ and a C/N ratio of 40 is assumed on day 200. This scenario serves as a reference setting upon which further scenarios are compared to and built.

Scenario 1Stat: Excluding particle movement

A key feature of the presented approach is the coupling of structure dynamics to C turnover and vice versa as already demonstrated in the studies of Rötzer et al. (2023); Zech et al. (2022b, 2024). To investigate the impact of the spatial dynamics in our setting, i.e., the rearrangement of particles, and the coupling of the C and N turnover to the evolution of the pore system, we compare the reference Scenario 1Full - where organic and mineral particles and aggregates relocate over time - to results simulated with a static geometry. For the latter, the initial aggregated state (see Section 2.1) is chosen.

2.4.2 Scenario 2: Impact of the quality and amount of exudates

Scenario 2CN: Impact of the quality of root exudates

To study the impact of exudate quality on the C and N cycles, three different C/N ratios are assumed for the single exudation pulse at day 200 which are compared to the reference Scenario 1Full ($C/N = 40$): $C/N = 10$, which is in the range reported for maize mucilage (e.g. Brax et al. (2020)), $C/N = 22$ (Niedeggen et al., 2024) and a C-rich situation is considered by setting $C/N = 100$. The total amount of C exuding is the same in these settings, only the corresponding amount of N varies in the sub-scenarios.



Table 4. Definition of model parameters and values used in the simulations with references, if available.

Symbol	Definition	Unit	Value	Reference
ν_{CS}	Maximum uptake rate	s^{-1}	$1.6 \cdot 10^{-5}$	Niedeggen et al. (2024)
k_{CS}	Half saturation constant	$\mu g C g^{-1}$	116	Niedeggen et al. (2024)
C_B^0	Total initial microbial biomass	$\mu g C g^{-1}$	158	Niedeggen et al. (2024)
C_B^{MIN}	Min. microbial biomass C cell conc.	$g C cm^{-3}$	0.0016	Zech et al. (2022a)
C_B^{MAX}	Max. microbial biomass C cell conc.	$g C cm^{-3}$	0.3168	Zech et al. (2022a)
N_B^0	Initial microbial biomass CAM cells	-	216	Zech et al. (2022a)
$(C/N)_S^0$	Initial C/N ratio of the substrates in the soil solution	-	10	
C_S^0	Initial C concentration in the soil solution	$g C cm^{-3}$	10^{-5}	Langeveld et al. (2020)
C^{Ex}	Min. C concentration creating a gluing spot	$g C cm^{-3}$	10^{-2}	
C_{POM}^0	Initial POM cell concentration	$g C cm^{-3}$	0.48	Zech et al. (2022b)
C_{Nec}^0	Initial necromass cell concentration	$g C cm^{-3}$	0.3168	
k_{POM}	Basal POM decay rate constant	d^{-1}	0.0096	Zech et al. (2022b)
k_{MN}	Basal necromass decay rate constant	d^{-1}	0.0096	
C_B^k	Microbially mediated factor, modifying rates k_{POM} & k_{MN}	$g C cm^{-3}$	6.3360	
D_S	Diffusion-dispersion coefficient of dissolved C and N	$cm^2 s^{-1}$	$1.94 \cdot 10^{-10}$	Kaiser et al. (2014)
$(C/N)_B$	Constant C/N ratio of the microbial biomass	-	10	
$(C/N)_{Nec}$	Constant C/N ratio of the microbial Necromass	-	10	
$(C/N)_{POM}$	Constant C/N ratio of POM	-	100	
$(C/N)_{Root}$	Constant C/N ratio of root exudates	-	40	Miao et al. (2020)
k_G	Respiration rate constant for growth and enzyme production		0.26	Kaiser et al. (2014)
k_M	Maintenance respiration rate constant	s^{-1}	$0.3 \cdot 10^{-6}$	Niedeggen et al. (2024)
k_{Nec}	Rate constant for microbial biomass turned into necromass	s^{-1}	$1.6 \cdot 10^{-6}$	Cochran et al. (1988)
k_{leak}	Rate constant for N leakage	h^{-1}	0.0001	Kaiser et al. (2014)
R_{root}	Exudation rate of C by root per time and root surface area	$\frac{\mu g C}{h \cdot cm^2}$	3.6	Santangeli et al. (2024)



Scenario 3Pulse: Impact of the amount of root exudates

255 In natural soils, multiple fine roots can influence the same soil region within a growth period. To mimic such a situation we establish 10 successive one-day-pulses of root exudation with $C/N = 40$ as they were described in Scenario 1Full. Within a time span of 50 days and at a frequency of 5 days, one day of exudate pulse is followed by 4 days without. The first pulse starts at time $t = 200$ d.

3 Results

260 The simulation results are analyzed in terms of the dynamics of the distinct C and N pools (recall Fig. 1). We distinguish the initially present necromass from the one created in the course of the simulation and quantify the relative share of C and N (in %) in the various pools. We also specify the fluxes between them, as defined in Section 2 and depicted in Fig. 1. The Carbon Use Efficiency (CUE) of the microbes is evaluated as ratio between uptake and assimilated carbon (Manzoni et al., 2018):

$$CUE = 1 - \frac{R_O + R_M + R_G}{f_U}. \quad (11)$$

265 Considering the estimated total dry mass soil in the domain, see Section 2.3, the soil C [$\text{mgNg}^{-1}\text{soil}$] and N [$\text{mgNg}^{-1}\text{soil}$] concentrations are calculated for each pool in each timestep and their temporal evolution is depicted.

As indicators of potential hotspots, i.e regions of elevated microbial activity or increased concentrations of microbial remnants, driven by processes such as EPS production, root exudation, and exoenzyme-mediated decomposition of POM or necromass, we define the relative availability of dissolved C for the microbes: it is defined by the fraction term $\frac{C_S}{C_S + k_{C_S}} \in [0, 1)$ in the Michaelis-Menten kinetics (Table 2) which indicates portion of the maximum microbial growth can be attained.

3.1 Scenario 1: C and N dynamics under continuous POM input and short-term root exudation

We first depict the results of Scenario 1Full (cf. Section 2.4). The temporal development of C and N concentrations, their fluxes and distribution among the different pools (see Fig. 5) reveal three different phases: the initial state up to day 199 (first equilibrium), the time span under the influence of the exudate pulse (days 200 - 250), and the final phase from day 250 up to day 500 (second equilibrium).

Initial Phase: Day 1-199

Dynamics of C and N pools

280 The evolution of the C concentrations in the distinct pools are depicted in Fig. 5a and the respective shares are illustrated in Fig. 5e. We observe the increase of POM due to its continuous addition (dark green curve), the slow degradation of initially present necromass (dark brown) and the generation of new necromass over time (light brown) due to the death of living biomass. Departing from the initial state the C concentrations in the soil solution (red curve), the total necromass (orange curve) and – after approximately 75 days – also the microbial biomass (light green curve) tend to dynamic equilibria. The N concentrations show the corresponding behavior (Fig. 5b and Fig. 5f).

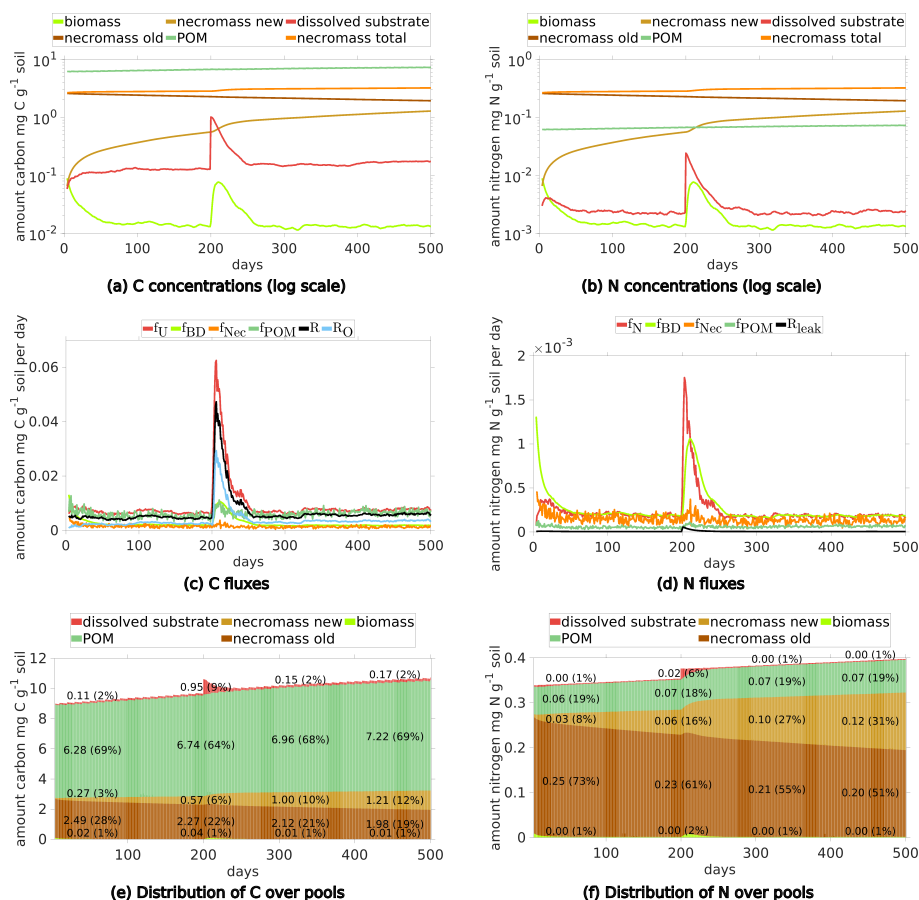


Figure 5. Temporal evolution of C and N concentrations (a and b) and C and N fluxes between different pools (c and d), and the relative distribution of total C and N in different pools (e and f) for scenario 1Full over 500 days. Continuous input of POM, one exudate pulse (C/N 40) after 200 days. Remember the definition the fluxes and pools, cp. Fig. 1. The concrete values of the distribution in e) and f) are depicted at day 50, 200, 325 and 450

Under the impact of continuous POM input and turnover of OM, the size of the microbial biomass reaches its dynamic equilibrium of 0.015 mgCg⁻¹soil after around 75 days. This corresponds to 10% of the initial amount of biomass. As such the biomass makes up less then 1 % of the total SOC (see Fig. 5e), which is common for agricultural soils Wei et al. (2022).

Likewise, necromass decomposition and its simultaneous formation through the decay of microbial biomass result in a dynamic equilibrium of approximately 2.8 mgCg⁻¹soil, which is within the range of the initial amount of necromass. As illustrated in Fig. 5e, this sums up to a fraction of around 30 % of SOC, but around 80 % of total N (Fig. 5f) (Zhang et al. (2022) e.g. report 40% of SOC).



After a short initial increase of the dissolved C and N concentrations (cf. the red curves in Fig. 5a and 5b), they stabilize at approximately $0.12 \text{ mgCg}^{-1}\text{soil}$ and $0.0025 \text{ mgNg}^{-1}\text{soil}$ after 25 days, which is in the range of lab experiments conducted by Cookson et al. (2007).

This corresponds to a C/N ratio between 40 and 55 in the soil solution.

295 **Fluxes** The dynamic equilibria of the state variables are also evident from the balancing fluxes as depicted in Figures 5c and 5d. Up to the simulation time of 200 days the fluxes are fluctuating around their individual equilibria, where the decomposition rates of POM f_{POM} and necromass f_{Nec} exhibit the strongest fluctuations.

The similar magnitude of the influx i.e. the turnover of biomass into necromass f_{BD} and the outflux, i.e. the necromass decay f_{Nec} , both in range of $0.002 \text{ mgCg}^{-1}\text{day}^{-1}$, result in a dynamic equilibrium of the total necromass.

300 Another equilibrium can be identified based on the fluxes that govern the dynamics of microbial biomass. As an example, on day 199, the uptake of dissolved C ($f_U = 0.006 \text{ mgCg}^{-1}\text{day}^{-1}$) is equal to the sum of respiration ($R = 0.004 \text{ mgCg}^{-1}\text{day}^{-1}$) and microbial decay ($f_{BD} = 0.002 \text{ mgCg}^{-1}\text{day}^{-1}$).

In our simulation, the mean of the total respiration between day 75 and 199 is around $4.8 \text{ } \mu\text{gCO}_2\text{-Cg}^{-1}\text{day}^{-1}$, which aligns with Niedeggen et al. (2024), where a basal microbial activity of $0.18 \text{ } \mu\text{gCO}_2\text{-Cg}^{-1}\text{h}^{-1} = 4.32 \text{ } \mu\text{gCO}_2\text{-Cg}^{-1}\text{day}^{-1}$ was
305 measured. The respiration overflow (due to the imbalance of C/N for biomass) accounts for about 55 % of the total respiration.

Concerning N, the equilibrium in the system is reflected by the balance between assimilation f_N and loss through microbial decay f_{BD} , both amounting to $0.0002 \text{ mgNg}^{-1}\text{day}^{-1}$. The high N flux of f_{BD} during the first 20 days (light green, Fig. 5d) corresponds to the high initial degradation of biomass.

Influence of exudate pulse: days 200 - 250

310 **Pools:** After 200 days, a flush of root exudates with C/N 40 enters the system via the soil solution along a central line through the domain (see the description of Scenario 1Full in Sec. 2.4). As a consequence, a rapid increase of dissolved C by 789% to $1.01 \text{ mgCg}^{-1}\text{soil}$ (red peak in Fig. 5a) and of dissolved N by 1043 % to $0.024 \text{ mgNg}^{-1}\text{soil}$ (red peak in Fig. 5b) is observed. The nutrient availability results in a subsequent increase of the microbial biomass (light green in Figs. 5a and 5b), followed by a subsequent, modest increase of new microbial necromass (light brown). The peak of the microbial biomass with 0.076
315 $\text{mgCg}^{-1}\text{soil}$ is reached 11 days after the pulse. This corresponds to a growth of 506 % compared to the equilibrium value of $0.015 \text{ mgCg}^{-1}\text{soil}$. The N concentrations (see Fig. 5b) exhibit the same behavior as the C concentrations, but at different absolute levels.

Fluxes: Most evident are the high uptake fluxes of DOC and DON (red curves f_U and f_N in Figs. 5c and 5d, respectively)
320 which enable the subsequent microbial biomass growth. In the sequel the decomposition of microbial necromass with $f_{Nec} = 0.0037 \text{ mgCg}^{-1}\text{day}^{-1}$ and POM with $f_{POM}=0.011 \text{ mgCg}^{-1}\text{day}^{-1}$ into dissolved carbon have doubled, while the microbial biomass decay flux $f_{BD} = 0.01 \text{ mgCg}^{-1}\text{day}^{-1}$ is even 5.6 times higher than before the exudation pulse. The peak activity of the microbial biomass leads to a total respiration of $47 \text{ } \mu\text{gCO}_2\text{-Cg}^{-1}\text{day}^{-1}$ cf. black curve in Fig. 5c, compared to Niedeggen et al. (2024), who reported a peak of $2.54 \text{ } \mu\text{gCO}_2\text{-Cg}^{-1}\text{h}^{-1} = 60.96 \text{ } \mu\text{gCO}_2\text{-Cg}^{-1}\text{day}^{-1}$.

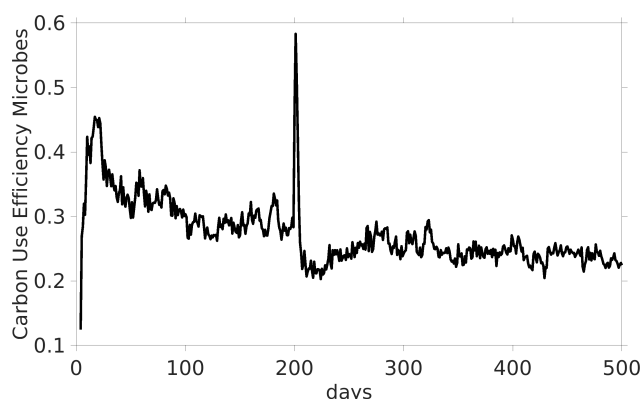


Figure 6. CUE of the reference scenario 1Full over time.

325 The increase of dissolved nitrogen by 1043 % is reflected by tenfold increase in N leakage.

The biomass dynamics is due to the (im-)balance of the N fluxes assimilation/mineralization f_N and biomass degradation f_{BD} : if $f_N \approx f_{BD}$ the microbial biomass stays in equilibrium (before the pulse). As the pulse enters the system $f_N \gg f_{BD}$ and the microbial biomass grows for 8 days. Then $f_N < f_{BD}$, and the microbial biomass degrades again.

330 **Second equilibrium: days 250 - 500**

The same equilibria for DOC and DON as before the exudate pulse are reached within 50 day, and with a delay of a few days also for the microbial biomass (see Figs. 5a and 5b). The main difference between pre- and post-pulse situation is the increase in microbial necromass from 2.65 to 3.23 $\text{mgCg}^{-1}\text{soil}$ (orange curves). POM rises slightly from 6.18 to 7.24 $\text{mgCg}^{-1}\text{soil}$, which each contributes 69 % to the SOC (see Fig. 5e). In total, this corresponds to a increase in TOC to 1.072 % of the soil's mass.

Also, all fluxes have returned to their equilibrium values (compare Fig. 5 before day 200 and after 260 days). Up to fluctuations, the equilibrium values remain constant until the end of the simulation time.

Finally, we present the temporal dynamics of the CUE of microbes. For the dynamic equilibrium between day 100 and 200 there is a CUE with a mean of 0.289. The CUE immediately reacts on the exudate pulse on day 200 and its peak is obtained on day 201 with a value of 0.583, see Fig. 6. As our model is capable of resolving localized processes, we illustrate the heterogeneous distribution of biomass and necromass in the domain along with the C availability exemplarily at three time steps in Fig. 7.

As the connectivity in the pore space essentially determines the local availability of dissolved C and N in the domain (Zech et al., 2022a), this results in locally variable microbial growth.

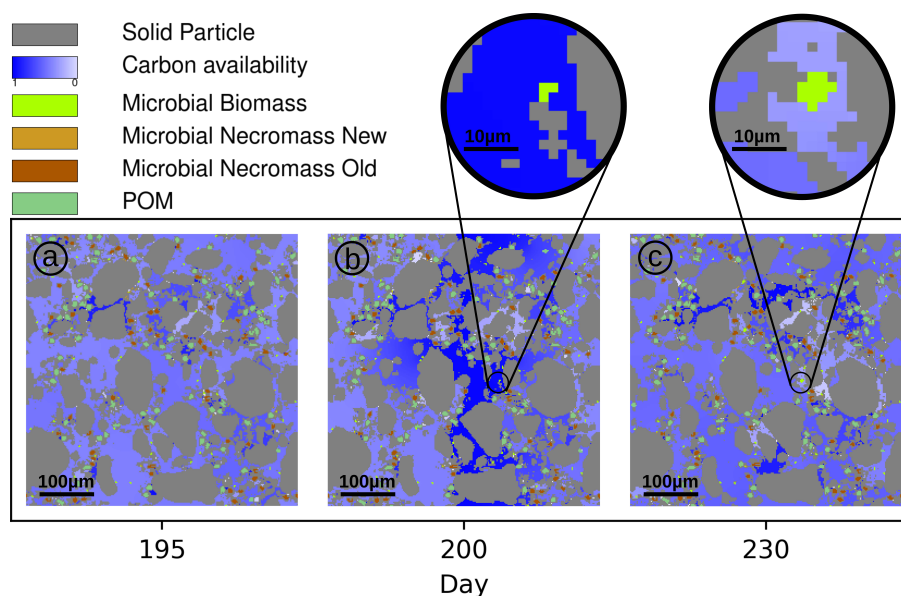


Figure 7. C availability (normalized; dark blue = high, light blue = low) in pore space in Scen. 1Full at days 195 (a), 200 (exudate pulse, b) and 230 (c). The lenses highlight a hot spot of biomass growth after the pulse.

If both substrate and consumer (microbes) are connected in small but isolated pores we observe that C is depleted, i.e. the availability (see definition above) is less than 0.4 (light blue regions in Fig. 7). In large, connected pore spaces, the amount of dissolved C available to the biomass is up to 0.7. Note that degradation of POM and microbial necromass also generates new DOC. On the other hand, in pores that are disconnected from living biomass, the C availability remains high, with values up to 1.0 (dark blue zones in Fig. 7). As we explicitly take into account the restructuring of the pore space, the connectivity may dynamically change.

At day 200 C-rich exudates enter the center of the domain (dark blue regions of Fig. 7b). Connected regions near the source of root exudation facilitate high C uptake rates by microbial biomass, leading to an increase in microbial growth (example: Figure 7, black circles). The microbial community grows in that hot spot under favorable conditions between day 200 (Fig. 7b) and 230 (Fig. 7c). In contrast, disconnectedness, i.e. physical barriers may prevent substrate exchange and so, new C-rich pores may persist after the exudate pulse because no consumers have access to them.

3.2 Scenario 1Stat: Impact of structure dynamics

We compare the reference scenario 1Full to scenario 1Stat, where the structural rearrangement is disregarded. The evolution of the C and N concentrations is depicted in Fig. 8, where we focus on the most dynamic phase from day 180 to 350. The most notable difference of the Scenario 1Stat is the increased concentrations of dissolved C and N (broken red lines in Fig. 8a and 8b), while the biomass concentrations exhibit a lower level (broken light green lines), e.g. C in microbial biomass tends to be 0.0044 mgCg⁻¹soil at day 199, which is only 30 % of the amount of C in biomass in Scenario 1Full. Ten days after the exudate

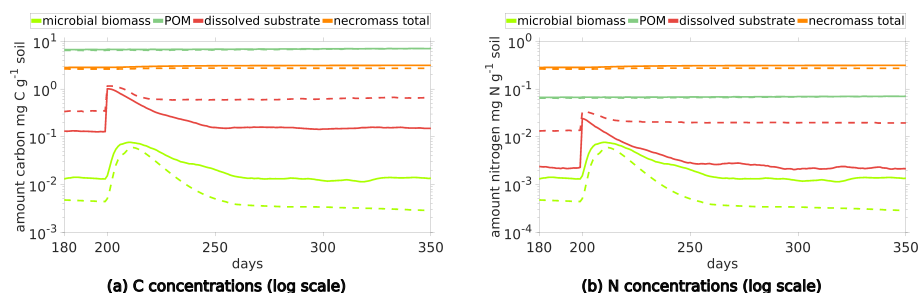


Figure 8. Temporal evolution of C and N concentrations in scenario 1Full (solid lines) and scenario 1Stat without particle movement (dashed lines).

pulse the microbial biomass peaks at 0.059 mgCg⁻¹soil, which is 23 % less compared to the scenario with particle movement. After 250 days, the effect of the pulse has vanished and dissolved N decreases slightly, while dissolved C continues to increase, but with less incline than before the pulse. In general the amount of dissolved carbon in scenario 1Stat is up to 5 times higher compared to scenario 1Full, dissolved nitrogen even up to 10 times higher.

Without movement of the particles TOC increases to 1.13 % of the soil's mass at the end of simulation. The difference of 0.62 mgCg⁻¹soil between both scenarios, 1Full and 1Stat, is a consequence of the lower CO₂ respiration in scenario 1Stat.

Furthermore it can be noted that the dissolved substrates and microbial biomass are fluctuating more in scenario 1Full, which is a direct consequence of the break-up of structures / bindings among the particles and thus changing connectivity between POM, necromass and biomass.

Figure 9 shows two snapshots at day 130 and 135 visualizing the C/N ratio in the soil solution to illustrate such an event. While the C/N distributions of the static geometries are almost identical (c and d, Scen. 1Stat), the C/N distribution changes in the reference scenario 1Full (a and b). The circled area points to the break-up of a structure that isolates two regions with high and low C/N at day 130. At day 135 the rearrangement of particles connected the two pore regions and the C and N concentrations equilibrated.

In general, the difference of the C/N ratios between unconnected pores is higher in Scenario 1Stat, leading to a pronounced patchiness. The amount of dissolved C and N in the pores is depending on the connectivity: In pores that are separated from spots with microbial biomass by physical barriers the dissolved C accumulates and exhibits a high potential carbon availability and C/N ratio. In contrast, in isolated pores containing microbial biomass, C and N are consumed, but without fresh C and N sources the C/N ratio in solution can increase while the uptake decreases until the biomass is not able to fulfill its basal activity and degrades.

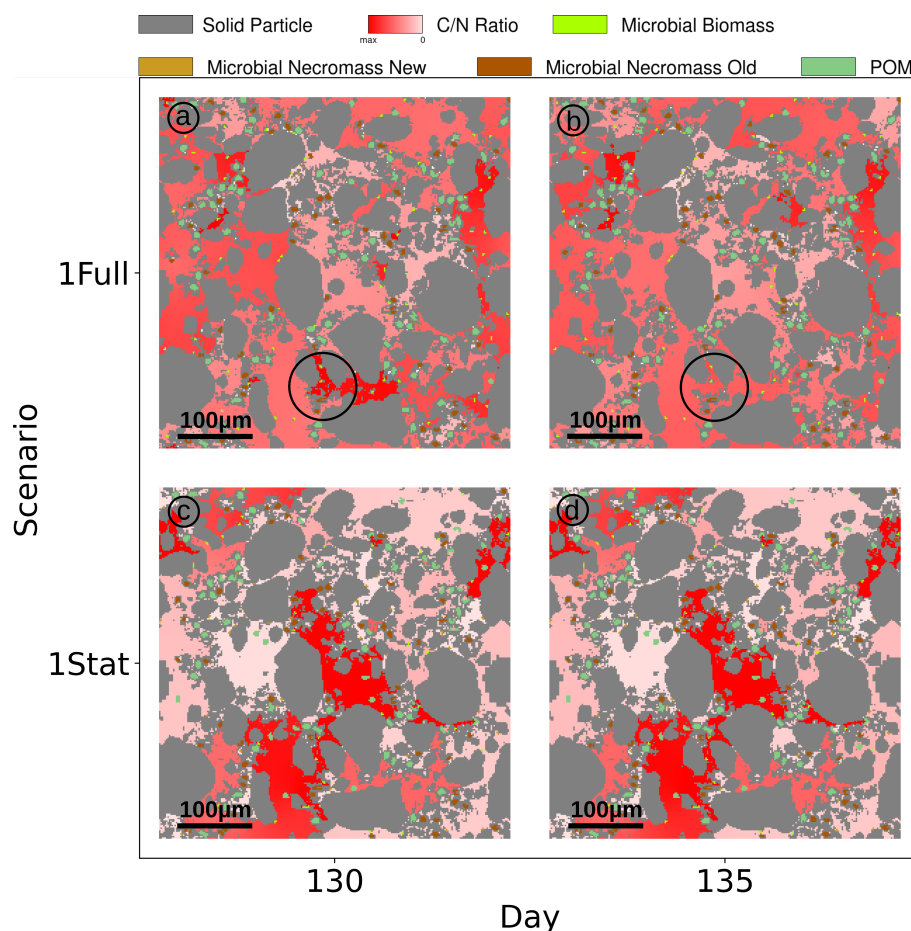


Figure 9. C/N ratio of soil solution at days 130 and 135 for the full reference scenario 1Full (a and b) and without particle dynamics 1Stat (c and d). Circle indicates where pores get connected by particle movement.

385 3.3 Scenario 2CN: Impact of the quality of root exudates

The alteration of the C/N ratio of the root exudates does not influence the evolution of the concentrations until the exudation pulse, so the same equilibria as in Scenario 1Full are established for all pools (see Fig. 10a and 10b). Small differences in the concentrations are due to the random break-up and placement of POM which may differ.

Investigating the impact of the exudation pulse, altering the C/N ratio of root exudates does not affect their carbon mass, resulting in similar dissolved C peaks (red lines in Fig. 10a). Dissolved N changes according to the given C/N ratio 10/22/40/100 in the sub-scenario (Fig. 10b), ranging from 0.01 mgNg⁻¹soil for a pulse of C/N 100, over 0.024 mgNg⁻¹soil (C/N 40) and 0.035 mgNg⁻¹soil (C/N 22) to 0.075 mgNg⁻¹soil for pulse of C/N 10. Higher amounts of dissolved N foster microbial growth, and accordingly C in biomass peaks are higher (light green curves in Fig. 10a) and the elevated level persists longer until the previous equilibrium state is reached after the pulse. For a pulse with C/N ratio of 10 it takes 125 days, while for a C/N ratio

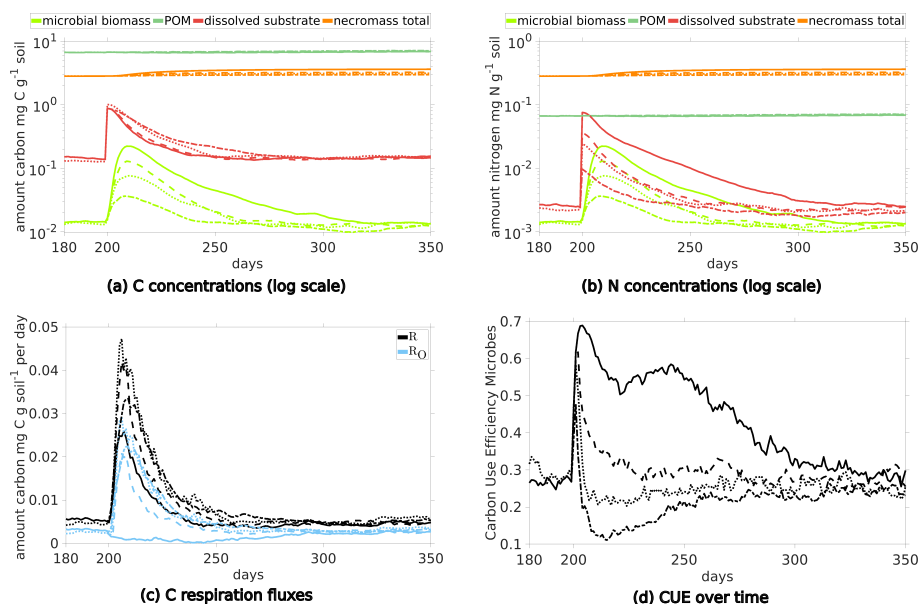


Figure 10. Temporal Evolution of C and N concentrations (a and b) and C respiration fluxes (c), and carbon use efficiency (d) for sub-scenarios comparing the impact of the quality of exudates: One pulse at day 200 with C/N ratios 10 (straight lines), 22 (dashed), 40 (dotted); Scen. 1Full) and 100 (dash-dotted).

of 100, less biomass is produced and the pre-pulse level is attained already after 50 days.

Studying the CUE (see Fig. 10d), it is evident that there is a different response for the exudates with C/N ratio of 10 and 22 compared to 40 and 100: All sub-scenarios exhibit a rapid increase in the CUE with peaks of 0.68 for C/N 10, 0.61 for C/N 22, 0.58 for C/N 40, and 0.47 for C/N 100. After the peak, CUE for C/N 40 and 100 briefly drops below their pre-pulse level of 0.3, recovering 50 and 100 days after the pulse, respectively. For C/N 22, the CUE drops to the pre-pulse state 30 days later, while for the optimal C/N 10 the microbial biomass maintains a CUE around 0.5 for a time span of almost 100 days.

Remark that a higher C/N ratio in root exudates corresponds to a prolonged elevated respiration (Fig. 10c dash-dotted and dotted lines), leading to a larger total loss in CO_2 .

The total respiration peak R caused by a root exudation pulse with C/N ratio 10 (solid lines) exhibits a value around $26 \mu\text{gCO}_2\text{-Cg}^{-1}\text{day}^{-1}$, the lowest among the different sub-scenarios. Strong differences are apparent comparing the respiration overflow of C due to the imbalance in C/N ratio in the biomass (light blue lines in Fig. 10c): While there is almost no C overflow when exudates had C/N ratio 10, there is C overflow in all other sub-scenarios.

The total amounts of added and lost C, and its distribution in the system at specific time points are summarized in Fig. 11. The microbial response to the different exudate qualities is indicated by the creation of new necromass during the simulations, as it results from the total amount of biomass ever existed. Comparison of day 200 with day 500 in Fig. 11 suggests that

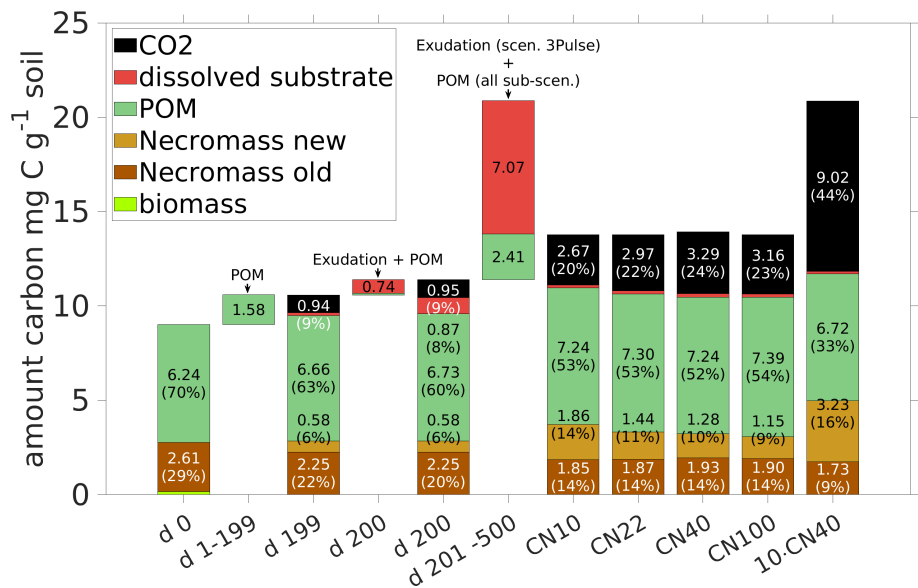


Figure 11. Total amount of C and distribution among the pools at time steps 0, 199, 200 (exudate pulse) and 500 is compared for different exudate qualities (C/N 10/22/40/100) in scenario 2CN and for the 10pulse-Scenario 3Pulse (C/N 40) (last bar). Additionally the amounts of C entering the system are displayed (POM and exudates) for the period 1-199 (POM, all scenarios), day 200 (exudates and POM, all scenarios), and days 201-500 (exudation in scenario 3Pulse, POM for all scenarios). For clarity, the values of pools contributing less than 2 % (biomass and dissolved substrate) are not shown.

higher C/N ratios in exudates lead to lower necromass accumulation but greater CO₂ release. From day 200 until day 500 in total 1.28 mgCg⁻¹soil necromass is generated from biomass if the pulse had a C/N ratio of 10, and simultaneously 1.72 mgC-CO₂ gsoil⁻¹ is released. On the other end, for a pulse with C/N 100 only 0.57 mgCg⁻¹soil necromass is generated, while 2.31 mgCg⁻¹soil is respired.

3.4 Scenario 3Pulse: Impact of the amount of root exudates

The ten pulses (at C/N 40) over 50 days increase dissolved C and N levels (red dotted curves in Figs. 12a and 12b). In response, the biomass also grows longer, and up to 0.31 mgCg⁻¹soil (dotted light green curve in Figs. 12a) compared to any scenario with one pulse (1Full or sub-scenario of CN2 with optimal ratio C/N 10). The maximum biomass concentration following a single pulse with C/N 10 is of similar height, while the biomass peak induced by a single pulse with C/N 40 is at most $\frac{1}{3}$ of the other maxima. Consequently, more necromass accumulates, resulting in a higher TOC of 1.18 % by the end of the simulation. Notably, ten pulses result in a necromass of 3.23 mgCg⁻¹soil and a CO₂ production of 9.02 mgCO₂-Cg⁻¹soil in total, compared with one pulse with high quality with a necromass increase of 73 % to the cost of 238 % more respiration. This is illustrated by the comparison of the bar plots of the final states in Fig. 11 and also expressed by a low CUE (Fig. 12d).

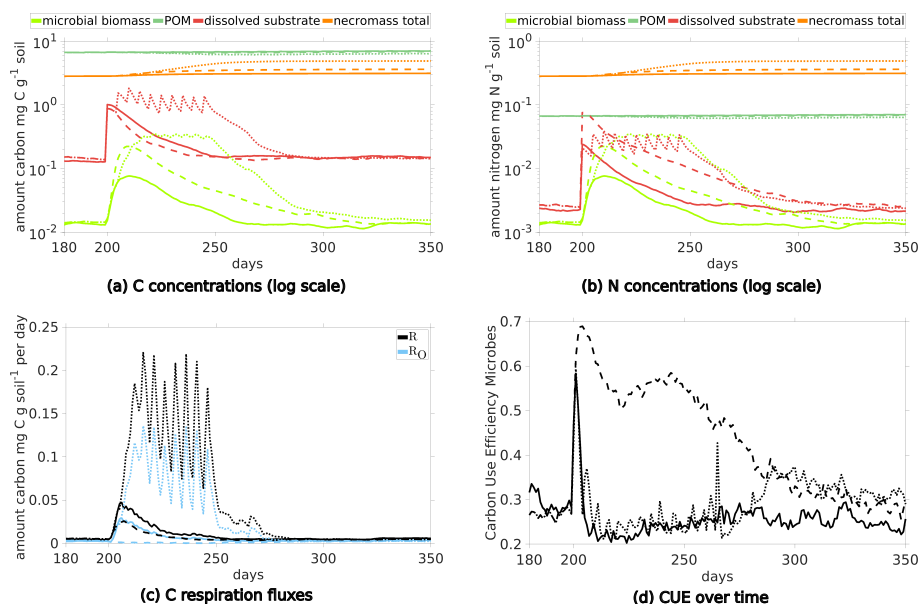


Figure 12. Temporal Evolution of C and N concentrations (a and b) and C respiration fluxes (c), and carbon use efficiency (d) for scenarios comparing the impact of the amount of exudates: Comparison of one pulse (at day 200) scenario with CN40 (straight lines, scenario 1Full) and CN10 (dashed, sub-scenario of 2CN) to 10 pulses with C/N 40 from day 200 to 250 (dotted).

425 Although the amount of C and N in the exudate pulses now is ten times higher than in the reference scenario 1Full, a high share is respired.

Comparing the distribution at day 500, we observe that the amount of C in the pools sums up to 56 %, while 44 % has been respired. For the single pulse with C/N 40, only 24 % of the C have been transformed to CO₂ while 76 % are still stored in the domain. We note that among all our presented scenarios, one pulse with C/N 10 produces least CO₂ (in total 2.67
430 mgCO₂-Cg⁻¹soil, 20 %) and stores second most C (summing up to 80 % and 11.1 mgCg⁻¹soil) .

4 Discussion

Discussing the results presented in Section 3 we highlight three key insights: the role of soil structure and local connectivity, nitrogen limitation, and on the spatial context of accumulation of microbial necromass.

4.1 The importance of heterogeneity and connectivity

435 Under the given model assumptions our study underlines that an evolving physical structure plays a critical role for C turnover (Lucas et al., 2021). In particular, the fully dynamic forward simulations (1Full) including a mechanistic view on OM turnover without parameter fitting reveal the emergence of equilibria in the different C and N pools, as observed in natural soils (Lehmann and Kleber, 2015).



The connectivity of nutrient sources and consumers, provided by the soil solution, alters due to the dynamic rearrangement of pores and solid matter, defining microenvironments with distinct locally varying concentrations of dissolved organic matter and biomass (see Fig. 7 and 9).

Microbial biomass patches residing in pores directly connected to sources of root exudates or decaying organic material benefit from favorable conditions promoting localized growth. In contrast, microbial biomass in isolated regions without direct access to substrate—due to limited pore connectivity—experience resource limitation and consequently decay.

Thus, local (dis-)connectivity can inhibit access and consumption, resulting in the observed increase by 5 to 10 times of dissolved C and N, and simultaneously 30 % lower biomass concentrations in the static scenario 1Stat (Fig. 8). This effect is also evident for root exudates as their deposition is restricted by soil structure and thus accumulates near the root within an average distance of about 100 μm , see Fig. 7 day 200. This localized enrichment of C and N creates hotspots for biomass growth, highlighting how physical barriers shape nutrient distribution and microbial growth in the rhizosphere. Furthermore, the addition of labile C and growth of biomass initiates gluing spot evolution and thus has an effect on structure dynamics, resulting in a feedback loop between microbial biomass and soil properties (Philippot et al., 2024). Already in Zech et al. (2022b) it was pointed out that higher organic matter input (there in terms of POM) led to larger aggregates which was mainly due to increased amount of organo-mineral associations.

These local concentration differences are highly sensitive to structural disruptions. Continuous (dis-)aggregation processes trigger the (de-)occlusion of POM and necromass, leading to constantly changing localized availability of C and N due to their accessibility-dependent decomposition (see in Fig. 7 and 9). Simultaneously, the restructuring of the soil matrix causes continuous changes in pore connectivity. Together, these two local factors - substrate accessibility and connectivity - govern the spatial distribution and availability of substrates, resulting in a dynamic rather than static equilibrium of microbial biomass (Fig. 8). Static geometries lack an important aspect mimicking the dynamic equilibria encountered in natural soils (Lehmann and Kleber, 2015).

4.2 Nitrogen limitation outweighs carbon surplus

By extending the C turnover model respecting the N requirements of the microbes we were able to trace and quantify the interplay between the C and N concentrations and biomass development in the context of a loamy, agricultural soil (Vetterlein et al., 2021). The turnover of living microbial biomass into necromass was also included and provides an important pool of substrates (Liang et al., 2019), along with POM. The system behavior during root exudation of maize plants was studied.

All simulation scenarios confirm that microbial biomass dynamically adapts its size to the locally available nutrients (dissolved C and N). The C uptake rate is primarily driven by the amount of DOC while the maximum possible growth rate is constrained by the amount of the available N (Schimel and Weintraub, 2003). We observed that, when a continuous supply of dissolved substrate is maintained —via successive degradation of necromass and POM, along with the regular addition of fresh POM particles— the microbial biomass stabilizes at a certain level where C and N uptake from the soil solution becomes constant. The stabilization occurs independently of the current biomass' size as e.g. defined by the initial conditions or short-term fluc-



tuations due to root exudation. This underlines that microbial biomass dynamics are primarily driven by the current availability of C and N (Kaiser et al., 2014).

475 The persistent overflow respiration of microbial biomass observed across all scenarios suggests an imbalance for the microbial metabolism due to a lack of N. N losses from biomass through its transformation into necromass can only be compensated by N uptake to a specific extent, underlined by the dynamic equilibria of fluxes. In contrast, C losses due to maintenance respiration and necromass formation can always be rebalanced, the microbial biomass responds with elevated respiration rates to the surplus of C, and to counterbalance the nitrogen limitation hampering growth. As a substantial portion of the organic matter
480 present in the system exists as POM, with a C/N ratio ten times higher than that required by microbial biomass, this results in the surplus of C. In scenarios, where the C/N ratio of the exudates is closer to 10, biomass growth was stimulated. Under such conditions, overflow respiration was markedly reduced, resulting in an considerably increased CUE, see also Manzoni (2017).

As such, in the scenario with the lowest C/N ratio of 10 (2CN), a temporary nitrogen saturation suggests that microbial biomass can adapt to root exudation dynamics, achieving the highest CUE.

485 A comparison of peak respiration following root exudate input under different C/N ratios suggests that no direct conclusions can be drawn regarding C assimilation or CUE based solely on these peak values. A peak respiration of $28 \pm 2.5 \mu\text{gCO}_2\text{-Cg}^{-1}\text{day}^{-1}$ could be observed in two distinct of our simulation scenarios: 2CN sub-scenario with C/N ratio 100 involves a smaller microbial biomass (peak of $0.036 \text{ mgCg}^{-1}\text{soil}$) exhibiting a high overflow respiration of $0.0224 \text{ mgC-CO}_2 \text{ gsoil}^{-1}\text{day}^{-1}$, in contrast sub-scenario with C/N ratio 10 exhibits a larger biomass (peak of $0.22 \text{ mgCg}^{-1}\text{soil}$), while maintaining a low overflow respiration of $0.001 \text{ mgC-CO}_2 \text{ gsoil}^{-1}\text{day}^{-1}$. Efficient C utilization becomes apparent only over longer time frames, during which
490 more necromass accumulates. This finding suggests that comparison between the initial and final system states provides more meaningful insights into microbial CUE and nitrogen limitation than a single-point respiration peak measurement. A single one-day exudation pulse with a favorable C/N of 10 elevated CUE more than 100 days, while 10 peaks with C/N 40 (3Pulse) could not increase CUE (Fig. 12).

495 **4.3 Long-term C and N storage by means of microbial necromass**

Due to the low N content of POM, the N required for biomass growth is primarily recycled from necromass. As a result, the overall accumulation of necromass stagnates or progresses only slowly. However, surplus C from POM efficiently supports microbial respiration, allowing the microbial community to maintain activity even under N-limited conditions.

During short term root exudation pulses, additional N becomes available and is rapidly taken up by the microbial biomass.

500 A portion of this biomass subsequently decays into necromass, effectively transforming readily available organic material into a more stable, slowly degradable form. Consequently, necromass accumulates and is incorporated into the soil matrix, contributing to long-term C and N storage (Buckeridge et al., 2022). This dynamics results in a net increase in necromass following the exudation event, with highest amounts of resulting necromass for lowest C/N ratios of the exudates (14 % of TOC for C/N ratio 10).

505 In our virtual loamy soils, POM and necromass can become physically occluded within fine soil particles. Without structural disruption, much of this organic matter remains protected from microbial decomposition, extending its residence time in the



soil (cf. Zech et al. (2022b)).

This is supported by the results of Scenario 1Stat (without structure dynamics), where both microbial biomass and respiration are significantly reduced compared to the scenario 1Full.

510 5 Conclusions

The presented approach provides novel insights into the microbially mediated processes that govern the spatial and temporal dynamics of carbon and nitrogen in soils at the micrometer scale. Rather than relying on models including fitted parameters, we employed forward simulations based on assumptions and parameter values directly derived from field and laboratory studies, allowing us to investigate the underlying drivers of observed phenomena in a mechanistic and transferable way. The results are coherent with experimental findings and support that the model incorporates the relevant mechanisms appropriately. Thus simulations allow us to analyze the distinct responses of the soil-microbe-carbon system to the input of slow-cycling or fast-turnover carbon.

As emphasized by Liang et al. (2019), next-generation field management needs to stimulate microbial biomass formation and necromass preservation to maintain healthy soils. The presented simulations support that their respective growth depends predominantly on the favorable C/N ratio of exudates rather than on a large amount, i.e. exudate diversity has been shown to have an important impact on C stability. Conversely, in soils that are subject to frequent structural changes—such as agricultural croplands—microbial activity is notably higher due to the regular exposure and decomposition of previously inaccessible organic matter. This increased accessibility enhances microbial biomass formation, leading to higher elevated CO₂ emissions.

Future work could incorporate the dynamics of phosphorous compounds in a similar manner and thus study their importance for microbial growth in particular for agricultural soils (Kwabiah et al., 2003). Fungal bio/necromass constitutes another important C and N pool in soils (Liang et al., 2019) which can be treated in an analogous way. The effects of environmental factors as temperature or moisture variations are not investigated in this study, but could give rise to further extensions.

Code and data availability. The data and code supporting the findings of this study are not publicly available, but are available from the corresponding author upon reasonable request.

530 Appendix A: Supplementary Information

A1 Cellular Automaton Method

For the Cellular Automaton Method (CAM), a computational domain (here: $(500 \times 500 \times 1) \mu\text{m}^3$) is discretized into cells (here: voxels of size $2 \mu\text{m} \times 2 \mu\text{m} \times 1 \mu\text{m}$), which are the smallest entities of the model. A specific type or state is assigned to each of these cells: solid, pore, POM, biomass or necromass. Several cells may be combined to larger entities to mimic a primary mineral soil, POM or necromass particle or micro-aggregate. While soil, POM, necromass, and pore cells have exclusive states,



biomass cells may include both living bacteria and necromass due to the continuous turnover of living bacterial biomass into necromass. Pores as well as biomass cells may additionally contain dissolved substrates such as DOC and DON.

The basic principle of a CAM is the application of rules for the potential relocation or transformation of cells in every time step. In our model, this includes the movement of solid particles, the turnover of POM, the evolution of biomass and necromass, the exudation and distribution of dissolved C and N within certain CAM cells, and changing surface properties of the edges of the solid cells.

In every time step, primary mineral soil, POM and necromass particles as well as aggregates thereof, relocate to achieve the most attractive and stable configuration, for details see Ray et al. (2017); Rötzer et al. (2023); Rupp et al. (2018b); Zech et al. (2022b).

The range of potential movement per timestep (day) increases with decreasing size as small particles are more mobile than large particles/aggregates, see Zech et al. (2022b) for details. It varies from 10 μm for units smaller than 100 μm^2 over 2 μm for units between 1600 and 80,000 μm^2 , to larger aggregates, which remain immobile.

The attractiveness of spots within the potential range of movement and thus the stability of bindings is ranked by the contact types: The connection between permanent reactive surfaces of primary soil particles or aggregates is considered the least attractive, while organo-mineral bindings between POM or necromass particles and reactive surfaces of soil particles are more favorable, with smaller soil particles having a higher proportion of such reactive surfaces. The most attractive configuration is between biomass, POM or necromass particles and primary soil particles or aggregates via so-called gluing spots. The latter evolves at the edge between a solid cell and pore space cell, where the combined amount of C in the DOC and microbial biomass exceeds a threshold C^{EX} .

As outlined in Zech et al. (2022b) gluing spots undergo an aging process over time in the absence OM and turn into so called memory edges. Ultimately, this aging results in the loss of their surface-binding capacity and therefore stability of the respective binding.

Finally, a random break-up of bindings is included to implicitly represent mechanical stresses due to, e.g., agitation, root growth or earthworms. The more stable the binding, the less probable the break-up.

Besides these relocation mechanisms, we particularly include the transformation of biomass to our CAM. Although we assume that the living microbial biomass is immobile, its concentration can spread by growth to neighboring pore or biomass cells once the carrying capacity of a biomass cell, i.e., the maximum mass of living or dead bacterial cells it can contain, is reached. This concept is adapted from Portell et al. (2018); Zech et al. (2020) and scaled accordingly to our situation. The corresponding maximal C concentration C_B^{MAX} of the microbial biomass is reduced by half of the necromass concentration C_{Nec} in each CAM cell, assuming that the fragments of two dead bacterial cells, which no longer store water, occupy the space of one living cell.

The corresponding amount of N in the microbial biomass cells is redistributed likewise to maintain the C/N ratio.

On the other hand, if the microbial biomass concentration in a cell C_B falls below a certain threshold C_B^{MIN} , equivalent to the minimum C concentration required for a single bacterial cell according to Portell et al. (2018), the cell state transitions



570 from biomass to a exclusive pore state and the the remaining C and N are released as DOC and DON the soil solution.

If all microbial biomass is fully converted into necromass, the state is exclusive and the same rules and principles apply for each necromass cell as for POM, i.e. it can relocate and decompose (details in section 2.2 in main document).

A2 Initial aggregated state

575 To create an aggregated initial state after the initial random placement of particles, the solid building units, POM and necromass particles are relocated within our cellular automaton framework due to their mutual attraction over 500 steps (without reactions taking place, see also the work flow in Fig. 4).

A3 Implementation

The implementation and simulations of the combined CAM and differential equations model as described in Section A1 build
 580 upon Rupp et al. (2018b) and Zech et al. (2022b) and were realized in Matlab (Inc., 2024). The relocation of soil, POM, and necromass particles and aggregates are performed in the CAM framework within a time step of one day. For each such timestep and for each pore, biomass, POM and necromass CAM cell, all biochemical transformations and stoichiometry are calculated, followed by the redistribution of the microbial biomass. These processes are resolved in time with a timestep of one hour to account for the rapid reaction rates.

585 To account for the diffusive as well as reactive components of our model, a non-iterative splitting is employed. More precisely, first disregarding the diffusive parts in the system of equations, the corresponding system of ordinary differential equations is solved with explicit Runge-Kutta (ode45 in Matlab) for all state variables (compare Table 1 in main document). Thereafter, these solutions are employed as reactive term and the remaining diffusive problem, which is discretized by a Local Discontinuous Galerkin method and implicit Euler, is solved (Rupp et al., 2018a). The procedure is repeated until the larger
 590 time step of one day is reached.

A4 Calculation of C overflow

We assume that microbial biomass is strictly homeostatic, i.e. the C/N ratio of biomass $\left(\frac{C}{N}\right)_B$ is constant. This implies that the uptake of N from the soil solution must match the actual demand as defined by means of the C dynamics. It is calculated as follows:

$$595 \quad \left(\frac{N}{C}\right)_B \partial_t C_B = \left(\frac{N}{C}\right)_B (f_U - R_G - R_M - f_{BD}) = N_B \left((1 - k_G) \nu_{C_S} \frac{C_S}{C_S + k_{C_S}} - k_M - k_{BD} \right) \quad (A1)$$

Depending on whether this amount of required N is available in the soil liquid, the following cases are distinguished and the overflow rate R_O is defined:

$$R_O = \begin{cases} 0 & \text{if } \left(\frac{N}{C}\right)_B \partial_t C_B < N_S, \\ \partial_t C_B - \left(\frac{C}{N}\right)_B N_S & \text{if } \left(\frac{N}{C}\right)_B \partial_t C_B \geq N_S. \end{cases} \quad (A2)$$



In the latter case the overflow rate exactly balances the mismatch of demanded and available N, i.e. it accounts for a potential
600 N limitation.

Author contributions. Maximilian Rötzer designed and carried out the formal analysis, developed the methodology, implemented the software, prepared the visualizations, and drafted the original version of the manuscript. Nadja Ray and Alexander Prechtel contributed to the formal analysis, designed the methodological framework, secured funding, supervised the project, and drafted the original manuscript. Eva Lehndorff, Andrea Scheibe, and Henri Braunmiller contributed to the conceptual design of the study, provided expertise and served as key
605 contacts for questions related to the soil system, and critically revised and edited the manuscript. All authors have read and approved the final version of the manuscript.

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. We gratefully acknowledge financial support of the Deutsche Forschungsgemeinschaft (DFG) within the framework of the research unit 2179 “MAD Soil – Microaggregates: Formation and turnover of the structural building blocks of soils,” project 276972051,
610 and the priority programme 2089 (Rhizosphere spatiotemporal organization-a key to rhizosphere functions) under the project numbers 403660839 and 251268514. We also thank Ulla Roskopf and Stephan Peth (Institute of Earth System Sciences, University of Hannover, Germany) for providing data on the particle size distributions.



References

- Baumert, V. L., Vasilyeva, N. A., Vladimirov, A. A., Meier, I. C., Kögel-Knabner, I., and Mueller, C. W.: Root exudates induce soil macroaggregation facilitated by fungi in subsoil, *Frontiers in Environmental Science*, 6, 1–17, <https://doi.org/10.3389/fenvs.2018.00140>, 2018.
- Baveye, P. C., Otten, W., Kravchenko, A., Balseiro-Romero, M., Beckers, É., Chalhoub, M., Darnault, C., Eickhorst, T., Garnier, P., Hapca, S., Kiranyaz, S., Monga, O., Mueller, C. W., Nunan, N., Pot, V., Schlüter, S., Schmidt, H., and Vogel, H.-J.: Emergent Properties of Microbial Activity in Heterogeneous Soil Microenvironments: Different Research Approaches Are Slowly Converging, Yet Major Challenges Remain, *Frontiers in Microbiology*, 9, 1–48, <https://doi.org/10.3389/fmicb.2018.01929>, 2018.
- Brax, M., Buchmann, C., Kenngott, K., Schaumann, G. E., and Diehl, D.: Influence of the physico-chemical properties of root mucilage and model substances on the microstructural stability of sand, *Biogeochemistry*, 147, 35–52, <https://doi.org/10.1007/s10533-019-00626-w>, 2020.
- Bucka, F. B., Kölbl, A., Uteau, D., Peth, S., and Kögel-Knabner, I.: Organic matter input determines structure development and aggregate formation in artificial soils, *Geoderma*, 354, 113 881, <https://doi.org/10.1016/j.geoderma.2019.113881>, 2019.
- Buckeridge, K. M., Creamer, C., and Whitaker, J.: Deconstructing the microbial necromass continuum to inform soil carbon sequestration, *Functional Ecology*, 36, 1396–1410, 2022.
- Chandel, A. K., Jiang, L., and Luo, Y.: Microbial models for simulating soil carbon dynamics: A review, *Journal of Geophysical Research: Biogeosciences*, 128, e2023JG007 436, 2023.
- Cochran, V., Horton, K., and Cole, C.: An estimation of microbial death rate and limitations of N or C during wheat straw decomposition, *Soil Biology and Biochemistry*, 20, 293–298, 1988.
- Cookson, W., Osman, M., Marschner, P., Abaye, D., Clark, I., Murphy, D., Stockdale, E., and Watson, C.: Controls on soil nitrogen cycling and microbial community composition across land use and incubation temperature, *Soil Biology and Biochemistry*, 39, 744–756, 2007.
- Crawford, J. W., Verrall, S., and Young, I. M.: The origin and loss of fractal scaling in simulated soil aggregates, *European Journal of Soil Science*, 48, 643–650, <https://doi.org/10.1111/j.1365-2389.1997.tb00564.x>, 1997.
- Crawford, J. W., Deacon, L., Grinev, D., Harris, J. A., Ritz, K., Singh, B. K., and Young, I.: Microbial diversity affects self-organization of the soil - Microbe system with consequences for function, *Journal of the Royal Society Interface*, 9, 1302–1310, <https://doi.org/10.1098/rsif.2011.0679>, 2012.
- Dungait, J. A., Hopkins, D. W., Gregory, A. S., and Whitmore, A. P.: Soil organic matter turnover is governed by accessibility not recalcitrance, *Global Change Biology*, 18, 1781–1796, <https://doi.org/10.1111/j.1365-2486.2012.02665.x>, 2012.
- Endress, M.-G., Dehghani, F., Blagodatsky, S., Reitz, T., Schlüter, S., and Blagodatskaya, E.: Spatial substrate heterogeneity limits microbial growth as revealed by the joint experimental quantification and modeling of carbon and heat fluxes, *Soil Biology and Biochemistry*, 197, 109 509, 2024.
- Gregory, P. J., George, T. S., and Paterson, E.: New methods for new questions about rhizosphere/plant root interactions, *Plant and Soil*, 476, 699–712, <https://doi.org/10.1007/s11104-022-05437-x>, 2022.
- He, X., Abs, E., Allison, S. D., Tao, F., Huang, Y., Manzoni, S., Abramoff, R., Bruni, E., Bowring, S. P. K., Chakrawal, A., Ciais, P., Elsgaard, L., Friedlingstein, P., Georgiou, K., Hugelius, G., Holm, L. B., Li, W., Luo, Y., Marmasse, G., Nunan, N., Qiu, C., Sitch, S., Wang, Y.-P., and Goll, D. S.: Emerging multiscale insights on microbial carbon use efficiency in the land carbon cycle, *Nature Communications*, 15, 8010, <https://doi.org/10.1038/s41467-024-52160-5>, 2024.
- Inc., T. M.: MATLAB version: R2024a, <https://www.mathworks.com>, 2024.



- 650 Kaiser, C., Franklin, O., Dieckmann, U., and Richter, A.: Microbial community dynamics alleviate stoichiometric constraints during litter decay, *Ecology Letters*, 17, 680–690, 2014.
- Kleber, M., Sollins, P., and Sutton, R.: A conceptual model of organo-mineral interactions in soils: self-assembly of organic molecular fragments into zonal structures on mineral surfaces, *Biogeochemistry*, 85, 9–24, <https://doi.org/10.1007/s10533-007-9103-5>, 2007.
- Kwabiah, A. B., Palm, C. A., Stoskopf, N. C., and Voroney, R. P.: Response of soil microbial biomass dynamics to quality of plant materials with emphasis on P availability, *Soil Biology and Biochemistry*, 35, 207–216, [https://doi.org/10.1016/S0038-0717\(02\)00253-5](https://doi.org/10.1016/S0038-0717(02)00253-5), 2003.
- 655 Kästner, M., Miltner, A., Thiele-Bruhn, S., and Liang, C.: Microbial Necromass in Soils—Linking Microbes to Soil Processes and Carbon Turnover, <https://doi.org/10.3389/fenvs.2021.756378>, 2021.
- Langeveld, J., Bouwman, A. F., van Hoek, W. J., Vilmin, L., Beusen, A. H., Mogollón, J. M., and Middelburg, J. J.: Estimating dissolved carbon concentrations in global soils: a global database and model, *SN Applied Sciences*, 2, 1626, 2020.
- 660 Lehmann, J. and Kleber, M.: The contentious nature of soil organic matter, *Nature*, 528, 60–68, <https://doi.org/10.1038/nature16069>, 2015.
- Liang, C., Amelung, W., Lehmann, J., and Kästner, M.: Quantitative assessment of microbial necromass contribution to soil organic matter, *Global Change Biology*, 25, 3578–3590, <https://doi.org/10.1111/gcb.14781>, 2019.
- Lucas, M., Vetterlein, D., Vogel, H. J., and Schlüter, S.: Revealing pore connectivity across scales and resolutions with X-ray CT, *European Journal of Soil Science*, 72, 546–560, <https://doi.org/10.1111/ejss.12961>, 2021.
- 665 Manzoni, S.: Flexible Carbon-Use Efficiency across Litter Types and during Decomposition Partly Compensates Nutrient Imbalances—Results from Analytical Stoichiometric Models, *Frontiers in Microbiology*, 8, 661, <https://doi.org/10.3389/fmicb.2017.00661>, 2017.
- Manzoni, S. and Porporato, A.: Soil carbon and nitrogen mineralization: Theory and models across scales, *Soil Biology and Biochemistry*, 41, 1355–1379, 2009.
- 670 Manzoni, S., Taylor, P., Richter, A., Porporato, A., and Ågren, G. I.: Environmental and stoichiometric controls on microbial carbon-use efficiency in soils, *New Phytologist*, 196, 79–91, 2012.
- Manzoni, S., Čapek, P., Porada, P., Thurner, M., Winterdahl, M., Beer, C., Brüchert, V., Frouz, J., Herrmann, A. M., Lindahl, B. D., Lyon, S. W., Šantrůčková, H., Vico, G., and Way, D.: Reviews and syntheses: Carbon use efficiency from organisms to ecosystems – definitions, theories, and empirical evidence, *Biogeosciences*, 15, 5929–5949, <https://doi.org/10.5194/bg-15-5929-2018>, 2018.
- 675 Miao, Y., Lv, J., Huang, H., Cao, D., and Zhang, S.: Molecular characterization of root exudates using Fourier transform ion cyclotron resonance mass spectrometry, *Journal of Environmental Sciences*, 98, 22–30, <https://doi.org/10.1016/j.jes.2020.05.011>, 2020.
- Nazari, M., Bickel, S., Benard, P., Mason-Jones, K., Carminati, A., and Dippold, M. A.: Biogels in soils: plant mucilage as a biofilm matrix that shapes the rhizosphere microbial habitat, *Frontiers in Plant Science*, 12, 798 992, 2022.
- Niedeggen, D., Rüger, L., Oburger, E., Santangeli, M., Ahmed, M., Vetterlein, D., Blagodatsky, S., and Bonkowski, M.: Microbial utilisation of maize rhizodeposits applied to agricultural soil at a range of concentrations, *European Journal of Soil Science*, 75, 1–14, <https://doi.org/10.1111/ejss.13530>, 2024.
- 680 Philippot, L., Chenu, C., Kappler, A., Rillig, M. C., and Fierer, N.: The interplay between microbial communities and soil properties, *Nature Reviews Microbiology*, 22, 226–239, 2024.
- Portell, X., Pot, V., Garnier, P., Otten, W., and Baveye, P. C.: Microscale Heterogeneity of the Spatial Distribution of Organic Matter Can Promote Bacterial Biodiversity in Soils: Insights From Computer Simulations, *Frontiers in Microbiology*, 9, 1–16, <https://doi.org/10.3389/fmicb.2018.01583>, 2018.



- Pot, V., Portell, X., Otten, W., Garnier, P., Monga, O., and Baveye, P. C.: Accounting for soil architecture and microbial dynamics in microscale models: Current practices in soil science and the path ahead, *European Journal of Soil Science*, p. ejss.13142, <https://doi.org/10.1111/ejss.13142>, 2021.
- 690 Ray, N., Rupp, A., and Prechtel, A.: Discrete-continuum multiscale model for transport, biomass development and solid restructuring in porous media, *Advances in Water Resources*, 107, 393–404, <https://doi.org/10.1016/j.advwatres.2017.04.001>, 2017.
- Rötzer, M., Prechtel, A., and Ray, N.: Pore scale modeling of the mutual influence of roots and soil aggregation in the rhizosphere, *Frontiers in Soil Science*, 3, <https://doi.org/10.3389/fsoil.2023.1155889>, 2023.
- Rupp, A., Knabner, P., and Dawson, C.: A local discontinuous Galerkin scheme for Darcy flow with internal jumps, *Computational Geo-*
695 *sciences*, 22, 1149–1159, 2018a.
- Rupp, A., Totsche, K. U., Prechtel, A., and Ray, N.: Discrete-Continuum Multiphase Model for Structure Formation in Soils Including Electrostatic Effects, *Frontiers in Environmental Science*, 6, 1–13, <https://doi.org/10.3389/fenvs.2018.00096>, 2018b.
- Rupp, A., Guhra, T., Meier, A., Prechtel, A., Ritschel, T., Ray, N., and Totsche, K.: Application of a Cellular Automaton Method to Model the Structure Formation in Soils Under Saturated Conditions: A Mechanistic Approach, *Frontiers in Environmental Science*, 7,
700 <https://doi.org/10.3389/fenvs.2019.00170>, 2019.
- Santangeli, M., Steininger-Mairinger, T., Vetterlein, D., Hann, S., and Oburger, E.: Maize (*Zea mays* L.) root exudation profiles change in quality and quantity during plant development – A field study, *Plant Science*, 338, 111 896, <https://doi.org/10.1016/j.plantsci.2023.111896>, 2024.
- Schimel, J. P. and Weintraub, M. N.: The implications of exoenzyme activity on microbial carbon and nitrogen limitation in soil: a theoretical
705 model, *Soil Biology and Biochemistry*, 35, 549–563, 2003.
- Schweizer, S. A., Bucka, F. B., Graf-Rosenfellner, M., and Kögel-Knabner, I.: Soil microaggregate size composition and organic matter distribution as affected by clay content, *Geoderma*, 355, 113 901, <https://doi.org/10.1016/j.geoderma.2019.113901>, 2019.
- Simon, E., Guseva, K., Darcy, S., Alteio, L., Pjevac, P., Schmidt, H., Jenab, K., Ranits, C., and Kaiser, C.: Distinct microbial communities are linked to organic matter properties in millimetre-sized soil aggregates, *The ISME Journal*, 18, <https://doi.org/10.1093/ismejo/wrae156>,
710 2024.
- Sokol, N. W., Whalen, E. D., Jilling, A., Kallenbach, C., Pett-Ridge, J., and Georgiou, K.: Global distribution, formation and fate of mineral-associated soil organic matter under a changing climate: A trait-based perspective, *Functional Ecology*, 36, 1411–1429, <https://doi.org/10.1111/1365-2435.14040>, 2022.
- Totsche, K. U., Amelung, W., Gerzabek, M. H., Guggenberger, G., Klumpp, E., Knief, C., Lehdorff, E., Mikutta, R., Peth, S.,
715 Prechtel, A., Ray, N., and Kögel-Knabner, I.: Microaggregates in soils, *Journal of Plant Nutrition and Soil Science*, 181, 104–136, <https://doi.org/10.1002/jpln.201600451>, 2018.
- Varney, R. M., Chadburn, S. E., Burke, E. J., and Cox, P. M.: Evaluation of soil carbon simulation in CMIP6 Earth system models, *Biogeo-*
sciences, 19, 4671–4704, <https://doi.org/10.5194/bg-19-4671-2022>, 2022.
- Vetterlein, D., Lippold, E., Schreiter, S., Phalempin, M., Fahrenkamp, T., Hochholdinger, F., Marcon, C., Tarkka, M., Oburger, E., Ahmed,
720 M., Javaux, M., and Schlüter, S.: Experimental platforms for the investigation of spatiotemporal patterns in the rhizosphere—Laboratory and field scale, *Journal of Plant Nutrition and Soil Science*, 184, 35–50, <https://doi.org/10.1002/jpln.202000079>, 2021.
- Vogel, H. J., Balseiro-Romero, M., Kravchenko, A., Otten, W., Pot, V., Schlüter, S., Weller, U., and Baveye, P. C.: A holistic perspective on soil architecture is needed as a key to soil functions, *European Journal of Soil Science*, 73, 1–14, <https://doi.org/10.1111/ejss.13152>, 2022.



- Vogel, H.-J., Amelung, W., Baum, C., Bonkowski, M., Blagodatsky, S., Grosch, R., Herbst, M., Kiese, R., Koch, S., Kuhwald, M., König, S., Leinweber, P., Lennartz, B., Müller, C. W., Pagel, H., Rillig, M. C., Rüschhoff, J., Russell, D., Schnepf, A., Schulz, S., Siebers, N., Vetterlein, D., Wachendorf, C., Weller, U., and Wollschläger, U.: How to adequately represent biological processes in modeling multifunctionality of arable soils, *Biology and Fertility of Soils*, 60, 263–306, <https://doi.org/https://doi.org/10.1007/s00374-024-01802-3>, 2024.
- Wang, B., Brewer, P. E., Shugart, H. H., Lerdau, M. T., and Allison, S. D.: Soil aggregates as biogeochemical reactors and implications for soil–atmosphere exchange of greenhouse gases—A concept, *Global Change Biology*, 25, 373–385, <https://doi.org/10.1111/gcb.14515>, 2019.
- Wei, L., Ge, T., Zhu, Z., Ye, R., Peñuelas, J., Li, Y., Lynn, T. M., Jones, D. L., Wu, J., and Kuzyakov, Y.: Paddy soils have a much higher microbial biomass content than upland soils: A review of the origin, mechanisms, and drivers, *Agriculture, Ecosystems & Environment*, 326, 107 798, <https://doi.org/https://doi.org/10.1016/j.agee.2021.107798>, 2022.
- Witzgall, K., Vidal, A., Schubert, D. I., Höschen, C., Schweizer, S. A., Buegger, F., Pouteau, V., Chenu, C., and Mueller, C. W.: Particulate organic matter as a functional soil component for persistent soil organic carbon, *Nature Communications*, 12, 4115, <https://doi.org/10.1038/s41467-021-24192-8>, 2021.
- Zech, S., Dultz, S., Guggenberger, G., Prechtel, A., and Ray, N.: Microaggregation of goethite and illite evaluated by mechanistic modeling, *Applied Clay Science*, 198, 105 845, <https://doi.org/10.1016/j.clay.2020.105845>, 2020.
- Zech, S., Ritschel, T., Ray, N., Totsche, K. U., and Prechtel, A.: How water connectivity and substrate supply shape the turnover of organic matter—Insights from simulations at the scale of microaggregates, *Geoderma*, 405, 115 394, 2022a.
- Zech, S., Schweizer, S. A., Bucka, F. B., Ray, N., Kögel-Knabner, I., and Prechtel, A.: Explicit spatial modeling at the pore scale unravels the interplay of soil organic carbon storage and structure dynamics, *Global Change Biology*, 28, 4589–4604, 2022b.
- Zech, S., Prechtel, A., and Ray, N.: Coupling scales in process-based soil organic carbon modeling including dynamic aggregation, *Journal of Plant Nutrition and Soil Science*, 187, 130–142, <https://doi.org/10.1002/jpln.202300080>, 2024.
- Zechmeister-Boltenstern, S., Keiblinger, K. M., Mooshammer, M., Peñuelas, J., Richter, A., Sardans, J., and Wanek, W.: The application of ecological stoichiometry to plant–microbial–soil organic matter transformations, *Ecological Monographs*, 85, 133–155, <https://doi.org/https://doi.org/10.1890/14-0777.1>, 2015.
- Zhang, Y., Gao, Y., Zhang, Y., Huang, D., Li, X., Gregorich, E., McLaughlin, N., Zhang, X., Chen, X., Zhang, S., et al.: Effect of long-term tillage and cropping system on portion of fungal and bacterial necromass carbon in soil organic carbon, *Soil and Tillage Research*, 218, 105 307, 2022.