

Spatiotemporal dynamics of Sentinel-2 NDVI as indicators of bio-hydromorphological interactions: implications for river management

Yuexia Zhou¹, Yuji Toda¹, Runye Zhu²

¹Department of Civil and Environmental Engineering, Nagoya University, Nagoya, Japan

²Zhejiang Institute of Hydraulics and Estuary, Hangzhou, China.

Correspondence to: Yuexia Zhou (zhou@civil.nagoya-u.ac.jp)

Abstract: The Normalized Difference Vegetation Index (NDVI) can be effectively used for monitoring the spatial and temporal dynamics of riparian vegetation. However, quantitative and efficient evaluations of the links between NDVI and bio-hydromorphological processes remain limited, particularly in river management contexts where dense in-channel vegetation can obstruct flow and reduce conveyance capacity. Using 200 cloud-free Sentinel-2 images (2015–2024) covering a 20-km reach of the Chikuma River (Japan), we evaluated the utility of high temporal resolution NDVI and greenness index (defined as $NDVI > 0.2$) as a quantitative indicator of bio-hydromorphological dynamics and its implications for riverine management. The analysis focused on the relationships between NDVI dynamics, flood magnitude, relative elevation along lateral channel morphology, and seasonal vegetation variability within a frequently disturbed channel. The results show that NDVI fluctuations strongly correspond to flood disturbances at lower relative elevations, whereas vegetation at higher elevations remains relatively stable. Along cross-channel transects, the annual maximum greenness ratio was well described by a logistic model, with parameter values varying according to the flood magnitude in the preceding year. Annual greenness ratio exhibited clear seasonal patterns, showing a later-summer (August–September) greenness peak. These spatiotemporal and seasonal NDVI characteristics demonstrate the potential of Sentinel-2 imagery to operationalize both the “where” (priority zones defined by relative elevation) and the “when” (timing) of vegetation management, providing a transferable, remotely sensed basis for flood-risk mitigation in frequently disturbed riverine environments.

1 Introduction

Riparian zones serve as dynamic interfaces between terrestrial and aquatic ecosystems, playing critical roles in flood regulation, nutrient cycling, biodiversity preservation, and habitat provisioning (Nallaperuma and Asaeda, 2019). The ecological functions of riparian zones result from continuous interactions amongst vegetation dynamics, sediment transport, channel migration, and river hydrodynamics, a process often referred to as bio-hydromorphological feedback (Gurnell et al., 2012; Naiman et al., 2010). Riparian vegetation typically exhibits high sensitivity to hydrological fluctuations such as floods. The seed dispersal and seedling recruitment was highly affected by the inundation zones during the seed dispersal period, the disturbance after seed dispersal, respectively. The flood-induced vegetation mortality may occur following uprooting, burial, bending and prolonged inundation time (Edmaier et al., 2011; Merritt et al., 2010). The riparian vegetation species and distribution are correlated with elevation relative to water level, which is traditionally expressed through the framework of lateral dimension (Hughes, 1997), emphasizing interactions between the river channel, the floodplain, and adjacent terrestrial habitats. Amongst these interactions, lateral connectivity is especially critical, since it enables the redistribution of sediment, nutrients, seeds, and organic materials across the river corridor, thereby supporting vegetation diversity and enhancing ecosystem resilience (Boothroyd et al., 2021; Hughes, 1997; Modi et al., 2022; Ward et al., 2002). Riparian vegetation cover increases flow resistance inducing sediment deposition and driving "biogeomorphic succession" (Corenblit et al., 2007). The river channel pattern and bank strength are thought to have a tight relationship with the in-channel riparian vegetation (Candel et al., 2021). Dense vegetation growth within river channels can influence river hydraulics by increasing channel roughness, thereby

40 impeding flow, elevating water levels, and potentially exacerbating flood risks (Bradley et al., 2007; Gao et al., 2022). The
intense flood events, which occurred in various countries around the world, highlighted the importance of evaluating the
vegetation destruction processes, such as the wood debris can induce dangerous obstruction along the channel network
(Francalanci et al 2020; Iwasaki et al., 2016). The embankment collapse was identified following the wash-out of riparian
vegetation (Kyuka et al., 2020). Understanding the destruction zone and its relationship with the hydromorphological
45 parameters, such as flood magnitude and relative elevation above the water level is necessary. The flood may cause vegetation
destruction and provide opportunities for post-flood succession (recruitment and growth) on newly formed bare substrates
(Mahoney et al., 1998; Merritt et al., 2010). The restoration of vegetation after flood pulse is also important for river managers
to make the management strategy from the medium to long-term viewpoints (Corenblit et al., 2011). Vegetation responses to
floods can include short-lived regrowth phases, high-frequency monitoring is critical to distinguish transient disturbance
50 signals from seasonal phenology and to capture recovery trajectories that may occur over weeks to months. This is especially
important in gravel-bed and braided systems where bare sediment exposure, and rapid geomorphic change complicate
interpretation from sparse surveys (Balke et al., 2014; Džubáková et al., 2015).

With the increasing availability of large geospatial datasets, the use of remotely sensed information to monitor riverine bio-
hydromorphodynamics has gained momentum (Singh and Vyas, 2022). The Normalized Difference Vegetation Index (NDVI),
55 derived from satellite remote sensing, is widely used to quantify vegetation greenness, canopy density, productivity, and
phenology (Marchetti et al., 2016; Redowan and Kanan, 2012; Zeng et al., 2020). Owing to the improved availability and
revisit frequency of satellite platforms, NDVI has enabled numerous spatiotemporal analyses of riparian vegetation, allowing
researchers to link vegetation phenology and interannual variability to hydrological and climatic drivers (Betz et al., 2023;
Henriques et al., 2024; Zuo et al., 2022). Despite these advances, existing studies have focused predominantly on lowland
60 rivers and watershed scale (Hess et al., 2003; Townsend and Walsh, 2001), and absolute elevation is often used as an
explanatory variable for NDVI dynamics. However, elevation typically decreases longitudinally from upstream to downstream,
and this trend does not adequately represent in-channel vegetation patterns. Consequently, quantitative evaluations of the
coupled effects of flood magnitude, lateral connectivity, and NDVI dynamics remain limited, particularly for high-disturbance
systems (Herrmann et al., 2024; Marchetti et al., 2020). In addition, practical frameworks that translate satellite-derived
65 vegetation indicators into operational guidance for managing channels prone to vegetation overgrowth remain underdeveloped.
Addressing these gaps is especially important where woody encroachment increases hydraulic roughness and flood risk.
This study addresses three primary research questions using a decade of remote sensing data from the Chikuma River, a steep,
gravel-bed and high-disturbance system in Japan characterized by significant vegetation overgrowth:

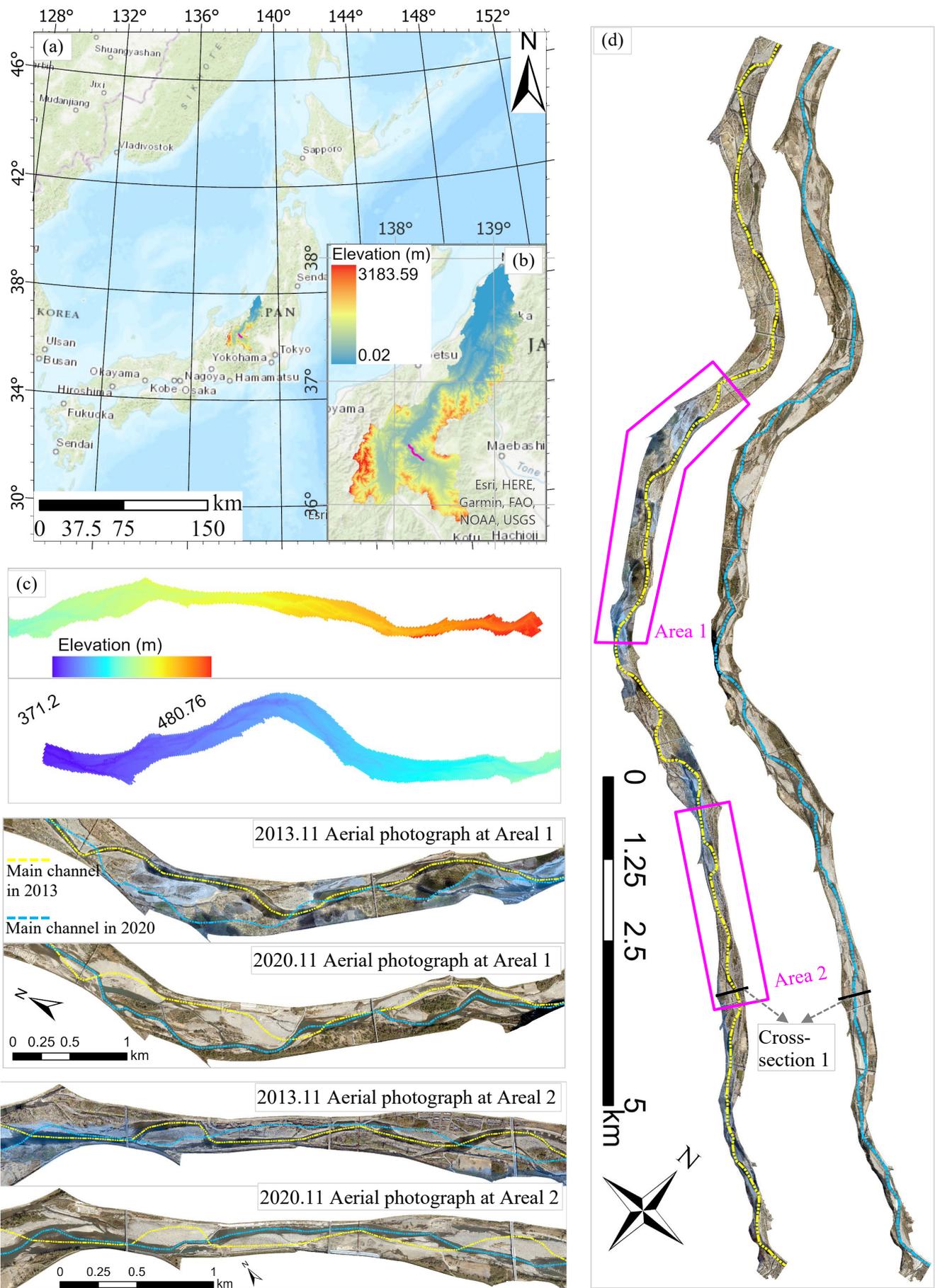
- 70 (1) How does the spatiotemporal Sentinel-2 NDVI respond to flood-disturbance and relative elevation above water level?
- (2) Can NDVI effectively indicate bio-hydromorphological interactions within frequent-disturbance system?
- (3) How can satellite-derived vegetation metrics be translated into practical strategies for river management?

By addressing these questions, we explore the potential of NDVI as a reliable indicator of bio-hydromorphological interactions
and assess its applicability for flood mitigation and vegetation control planning. The structure of the paper is as follows. We
first introduce datasets and analytical methods, including remote sensing, hydrological records, and topographic data
75 processing. We then evaluate spatial and temporal NDVI patterns in relation to flood magnitude and elevation relative to water
level, and we examine the seasonal variability of NDVI. Finally, we discuss the optimal time (“when”) and the spatial
prioritization (“where”) of riparian vegetation management actions.

2 Study area and data source

2.1 Study site

80 The study area for our research was the Chikuma River, located in Nagano Prefecture, Central Japan (Fig. 1). Our study



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Figure 1: An outline of the study area. (a) and (b) show the location of the target site in Japan (Esri, HERE, Garmin, FAO, NOAA, USGS) and the watershed (Digital Elevation Model (DEM) from the Geospatial Information (GSI) Authority of Japan), respectively; (c) shows the morphology (DEM from the GSI) of the target site; and (d) provide orthophotos from 2013 and 2020 (Copyright © Ministry of Land, Infrastructure, Transport and Tourism (MLIT) in Japan), respectively.

90 focused on a 20-kilometer river reach, extending from 109 km to 89 km upstream of the river mouth. Within this reach, the river has a slope of approximately 1/200 and a representative grain size (d_{60}) of 53 mm, classifying it as a steep, gravel-bed river. The 109–89 km section is treated as a single, coherent geomorphic segment because its channel characteristics are distinct from those of the adjacent reaches. The upstream section is characterized by lower vegetation coverage, while the downstream reach possesses a significantly gentler slope. Woody encroachment, primarily *Robinia pseudoacacia* and *Salix*, hinders flood management. In 2018, *Robinia pseudoacacia* alone accounted for approximately 23% of the total area, which is substantially higher than the coverage of bare gravel bars (around 11%). As a result, periodic, nature-oriented river maintenance practices, including vegetation cutting and river channel excavation, have been implemented for improving flood conveyance capacity. However, uncertainties remain regarding the optimal timing (“when”) and spatial prioritization (“where”) of vegetation cutting, posing challenges for river managers.

2.2 Data source

100 By considering the application of readily available high-frequency satellite images on grasping riparian vegetation dynamics, a time series of Sentinel-2 from 2015 to 2024 was collected. Topographic and hydrological data were additionally obtained to evaluate the dynamics of hydrology and channel lateral morphology.

2.2.1 Sentinel-2 images

105 For the study, we collected Sentinel-2 satellite imagery (Red, Green, Blue, and Near-Infrared (NIR) bands) at a spatial resolution of 10 m. Given that the minimum and average channel widths at the study reach are approximately 190 and 380 m, respectively, a spatial resolution of 10 m was sufficient to effectively distinguish between vegetation zones, bare gravel bars, and water surfaces. In total, 200 Sentinel-2 images were downloaded from Copernicus and processed to compute the NDVI. Imagery was selected based on cloud-free conditions to ensure reliable observation of spatiotemporal vegetation and channel dynamics over the period 2015–2024. Figure 2 presents Sentinel-2 imagery (2016–2024) for representative Area 2 in Figure 1. Moderate and pronounced vegetation mortality and channel shifts, are evident in the 2018 and 2020 images (highlighted by red circles in Fig. 2), corresponding to the second-largest and largest flood events during the study period, respectively.



Figure 2: Sentinel-2 imagery from 2016 to 2024 at representative Area 1

2.2.2 Hydrological data

115 Hydrological records, specifically hourly water level (Fig. 2) measurements during the study period, were obtained from the Water Information System maintained by Ministry of Land, Infrastructure, Transport and Tourism (MLIT), Japan. Data was acquired from the Kuiseke gauging station, located near the center of the study reach. These water level records allowed for the identification of flood timing and magnitude throughout the study period, thus providing critical inputs for investigating the relationship between flow disturbances and vegetation dynamics. During the 10-year study period, four relatively high-magnitude flood events were identified, ranked by severity as October 2019, October 2017, August 2021, and October 2018.

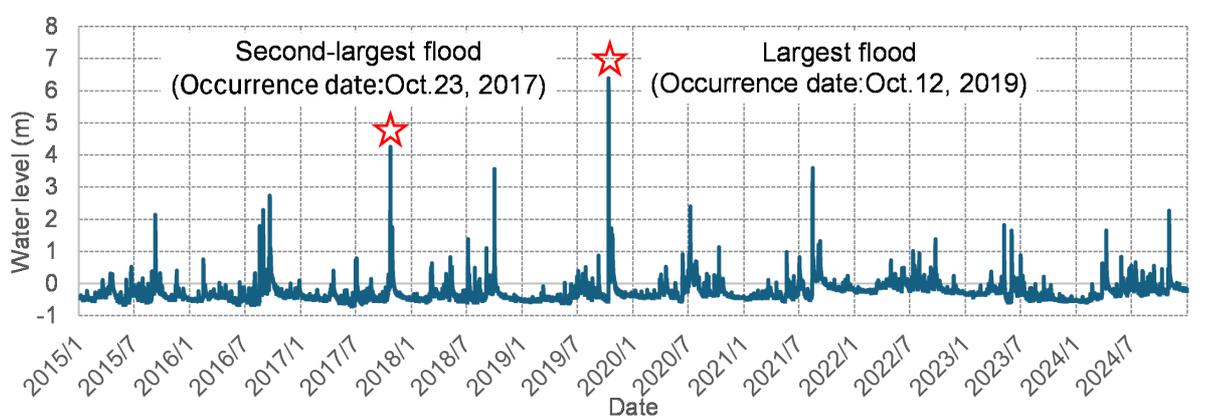


Figure 3: Temporal water level at the Kuisseke gauge station.

2.2.3 Topography data and vegetation height

125 Topographic information for areas above the water's surface was obtained from a laser profile (LP) survey, which provided
 the digital elevation model (DEM) and the digital surface model (DSM), with a resolution of 0.5 m. Since the LP sensor does
 not penetrate water, bathymetric data for submerged areas were derived from cross-sectional surveys conducted at 500 m
 intervals. LP surveys were conducted in October 2013, December 2019, and November 2020, while cross-sectional surveys
 were performed in 2013, 2017 and 2019. These datasets revealed dynamic changes in lateral channel morphology (Fig. 3) and
 130 shifts in the main channel from 2013 to 2019 (Fig. 1 (d)). Airborne Lidar Bathymetry (ALB) data surveyed during early
 December 2024, were also acquired, covering both submerged and non-submerged topographic regions. All the topographic
 datasets utilized in this study were provided by the Chikuma River Office of the MLIT, Japan, and are considered reliable and
 accurate for the purposes of this research.

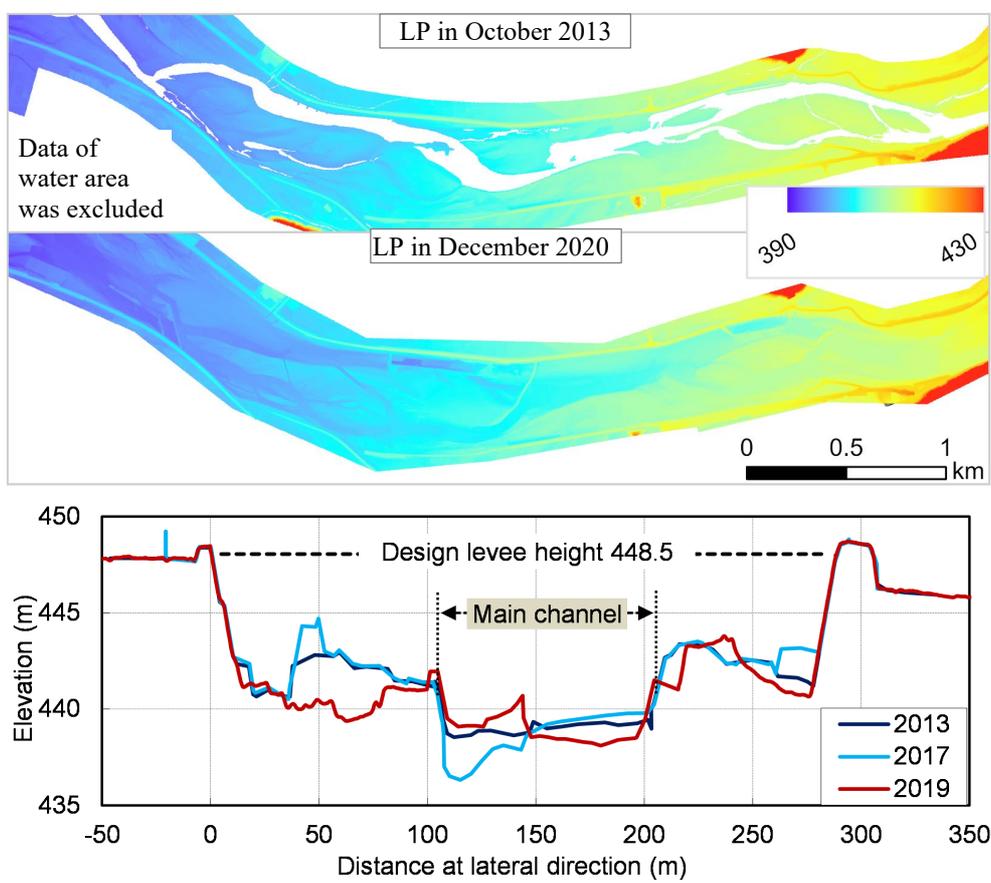


Figure 4: LP survey data at representative Area 1 and one cross-sectional profile of morphology at Cross-section 1.

2.3 The methodology for data analysis

2.3.1 The greenness ratio calculation based on NDVI data

NDVI values for each Sentinel-2 image were generated in ArcGIS Pro using Eq. (1). In previous studies, land cover has commonly been classified into three categories using NDVI thresholds: water (NDVI < 0), bare bars or sparse vegetation (NDVI = 0.0–0.2), and active vegetation (NDVI > 0.2) (Tucker, 1979; Drori et al., 2020). However, in riparian zones, NDVI values can drop below 0.2 during late autumn and winter due to seasonal senescence. As a result, we categorized areas with a NDVI ≤ 0 as water, a NDVI between 0 and 0.2 as bare bars or dormant (non-photosynthesizing) vegetation, and a NDVI > 0.2 as active photosynthesizing vegetation, which we refer to as “greenness” (Tucker, 1979). We specifically focused on NDVI > 0.2, vegetation greenness, for analyzing the interaction between vegetation dynamics, flood magnitude, and phenological characteristics, due to the following points. First, major floods at our study site typically occur during September or October, by which time vegetation has generally recovered from seasonal dormancy and is actively photosynthesizing (NDVI > 0.2). **During the floods period, vegetation cover is most fully developed and NDVI values above 0.2 reliably correspond to active green vegetation, making NDVI > 0.2 (Drori et al., 2020) a suitable threshold for estimating vegetation cover.** Second, phenological activity, inherently reflecting the growth and expansion of active vegetation, corresponding to NDVI values above 0.2. The greenness ratio is defined as the proportion of area with NDVI > 0.2 relative to total area (Eq. (2)). To quantify the temporal change of NDVI and the greenness ratio during the study period along the channel transect, we introduced the parameter $NDVI_{vary}$ and $Green_{vary}$, calculated using Eq. (3):

$$NDVI = NIR - Red / NIR + Red \quad (1)$$

$$Greenness\ ratio = N_{0.2} / N_{total} \quad (2)$$

$$\begin{cases} NDVI_{vary} = \sum_{y=2}^N \frac{Abs(\overline{NDVI}_y - \overline{NDVI}_{y-1}) / \overline{NDVI}_{y-1}}{y-1} \\ Green_{vary} = \sum_{y=2}^N \frac{Abs(\overline{Green}_y - \overline{Green}_{y-1}) / \overline{Green}_{y-1}}{y-1} \end{cases} \quad (3)$$

where, $N_{0.2}$ and N_{total} refer to the number of pixels within Sentinel-2 imagery where NDVI > 0.2 and the total number of pixels within the target area, respectively. \overline{NDVI}_y and \overline{NDVI}_{y-1} represent the average annual maximum NDVI values at each 0.5m relative elevation zone (detailed as described in detail in section 2.3.2) corresponding to y and $y - 1$ years, respectively. \overline{Green}_y and \overline{Green}_{y-1} represent the average annual maximum greenness ratio at each 0.5m relative elevation zone corresponding to y and $y - 1$ years, respectively. N is the total year during the study period. The term “Abs” represents the absolute value of the NDVI difference between two years.

2.3.2 **Relative elevation** above water level

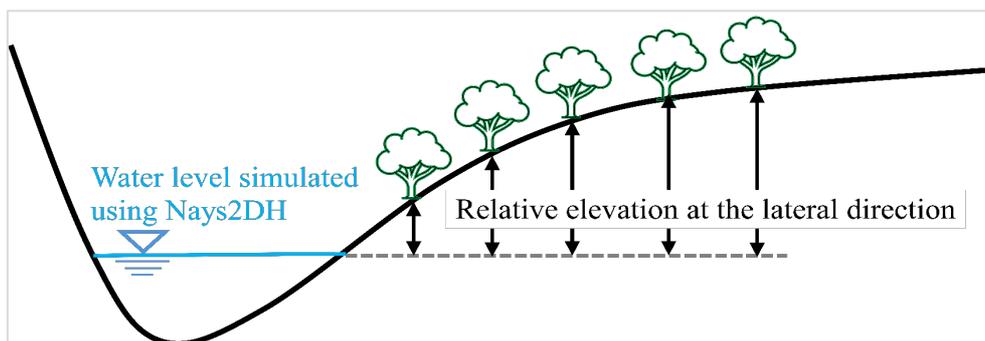
In this study, we focused on riparian vegetation located between levee banks within the river channel (**Fig. 5**). The conceptual framework of the lateral dimension, which emphasizes interactions between the river channel and riparian vegetation. The previous research (Zhu et al., 2023) suggests that relative elevation, defined as the height of floodplain landforms above the water’s edge, has a stronger influence on the vegetation distribution within levee banks, relative elevation is proposed as a useful index for characterizing lateral continuity and for examining the relationship between NDVI patterns and topographic characteristics within the river corridor. We specifically define relative elevation as the height above the 95th water level, indicating that the water’s surface is at or above this level approximately 95 days per year. Areas below this elevation are more frequently inundated by river flow, while areas above experience alternating wet and dry conditions that are more conducive for vegetation establishment and persistence. At our study site, riparian vegetation is predominantly distributed above the edge of the 95th flow inundation zone.



175 **Figure 5: A riparian vegetation lateral distribution (2024 ALB survey results, Copyright © MLIT in Japan).**

Riparian vegetation lateral distribution is often considered to be reset following major morphological change induced by large flood events (Death et al., 2015). The morphology change was identified after two significant floods occurred in 2017 and 2019 (Fig.3), prompting us to divide the study period into three intervals to better represent the evolving physical environment: Period 1 (2015–2017), Period 2 (2018–2019), and Period 3 (2020–2024).

180 To calculate relative elevations above water level for each period, we used a combination of topographic datasets. We specifically employed 2013 laser profiling (LP) and cross-sectional survey data for Period 1, 2013 LP and 2017 cross-sectional data for Period 2, and 2024 airborne lidar bathymetry (ALB) data for Period 3. Since floodplain morphology was assumed to largely remain unchanged, with most morphological adjustments occurring within the main channel (Fig. 4), the 2013 LP dataset was reused for both Periods 1 and 2. The 95th water level, which varies in response to morphological dynamics, was
 185 calculated for each analysis period using the two-dimensional hydraulic simulation model Nays2DH (Shimizu et al., 2020). Relative elevation was then derived by subtracting the simulated water level from topographic elevation (Fig. 6). For detailed analyses, resulting relative elevation values were classified into two schemes: one with 50 classes at 0.1 m intervals, and another with 10 classes at 0.5 m intervals, covering the range from 0 to 5 m.



190 **Figure 6: Conceptual diagram of relative elevation in the lateral direction.**

2.3.3 Seasonal dynamics of NDVI

In-channel riparian vegetation directly modifies hydraulic roughness and thus affects flow conveyance. Therefore, seasonal increases in vegetation greenness and biomass (as indicated by NDVI) can be interpreted as periods of potentially elevated flood risk, particularly when they coincide with the timing of major flood events. In-channel riparian vegetation directly
 195 modifies hydraulic roughness and thus affects flow conveyance; therefore, seasonal increases in vegetation greenness and

biomass (as indicated by NDVI) can be interpreted as periods of potentially elevated flood risk, particularly when they coincide with the timing of major flood events. In this study, we characterize the seasonal variability of NDVI as a proxy for seasonal roughness dynamics and discuss its implications for flood-risk management. Using time-series NDVI derived from Sentinel-2 imagery, we computed monthly mean NDVI for each year over a 10-year period. This aggregation reduces the influence of short-term fluctuations in water level and the noise associated with individual cloud-free scenes, allowing a clearer identification of recurring seasonal patterns that are relevant for operational planning (e.g., the timing of vegetation cutting prior to the typhoon season). To better understand the environmental drivers influencing these seasonal distribution patterns (Wang et al., 2024), we, additionally, collected and integrated temperature and precipitation data from the Ueda meteorological station into our analysis. Yearly averages of monthly mean temperature, monthly maximum temperature, and monthly accumulated precipitation were calculated for the period from 2015 to 2024.

3 Results

3.1 Spatial-temporal greenness ratio from 2015 to 2024

The NDVI dynamics at representative Area 1 (Fig.1) from 2016 to 2024 is shown as Figure 7. The area with NDVI values < 0.2 (highlighted by red circles in Fig. 7) decreased slightly and sharply following the 2017 and 2019 flood events, respectively, indicating flood-induced vegetation removal. Because vegetation destruction is influenced by both flood magnitude and inundation duration, and the typhoon-driven flood in this study was relatively short (<50 h), we used the 10 h accumulated hourly water level (a rolling sum of hourly water level over a 10 h window) as an integrated metric representing these two components. Figure 8 presents 10 hour accumulated water level alongside the greenness ratio at different relative elevations. The figure demonstrates the influence of the catastrophic flood on vegetation. Fluctuations in vegetation greenness were closely associated with both flood magnitude and relative elevation. The greenness ratio consistently declined following the occurrence of notable annual maximum floods, particularly those in 2017, 2018, 2019, and 2021. The severity of these floods, based on return periods, was as follows: 2019 (~50-year), 2017 (~10-year), 2021 (~5-year), and 2018 (~3-year). The observed reduction in greenness occurred across different lateral relative elevation ranges: 0–5 m (2019), 0–3 m (2017), 0–3 m (2021), and 0–2 m (2018). Although the lateral extent of greenness reduction during the 2021 and 2017 floods was similar, the magnitude of the reduction was greater in 2017. Overall, the magnitude of greenness declines mirrored flood severity: the most substantial reduction occurred during the 2019 flood, followed by the 2017 event, with smaller but still observable declines in 2021 and 2018, respectively.

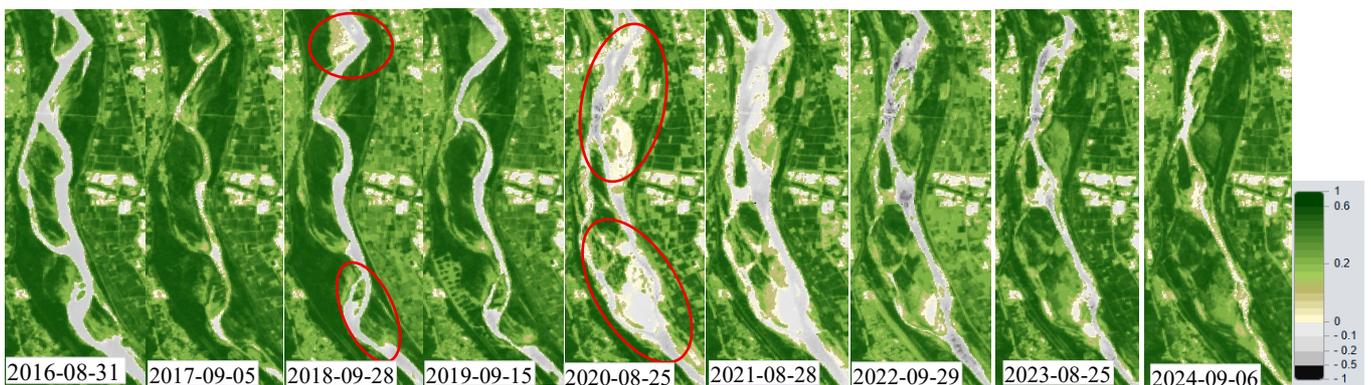


Figure 7: NDVI from 2016 to 2024 at the representative Area 1 (the vegetation mortality area after the largest flood and second-largest flood was marked with red polygon).

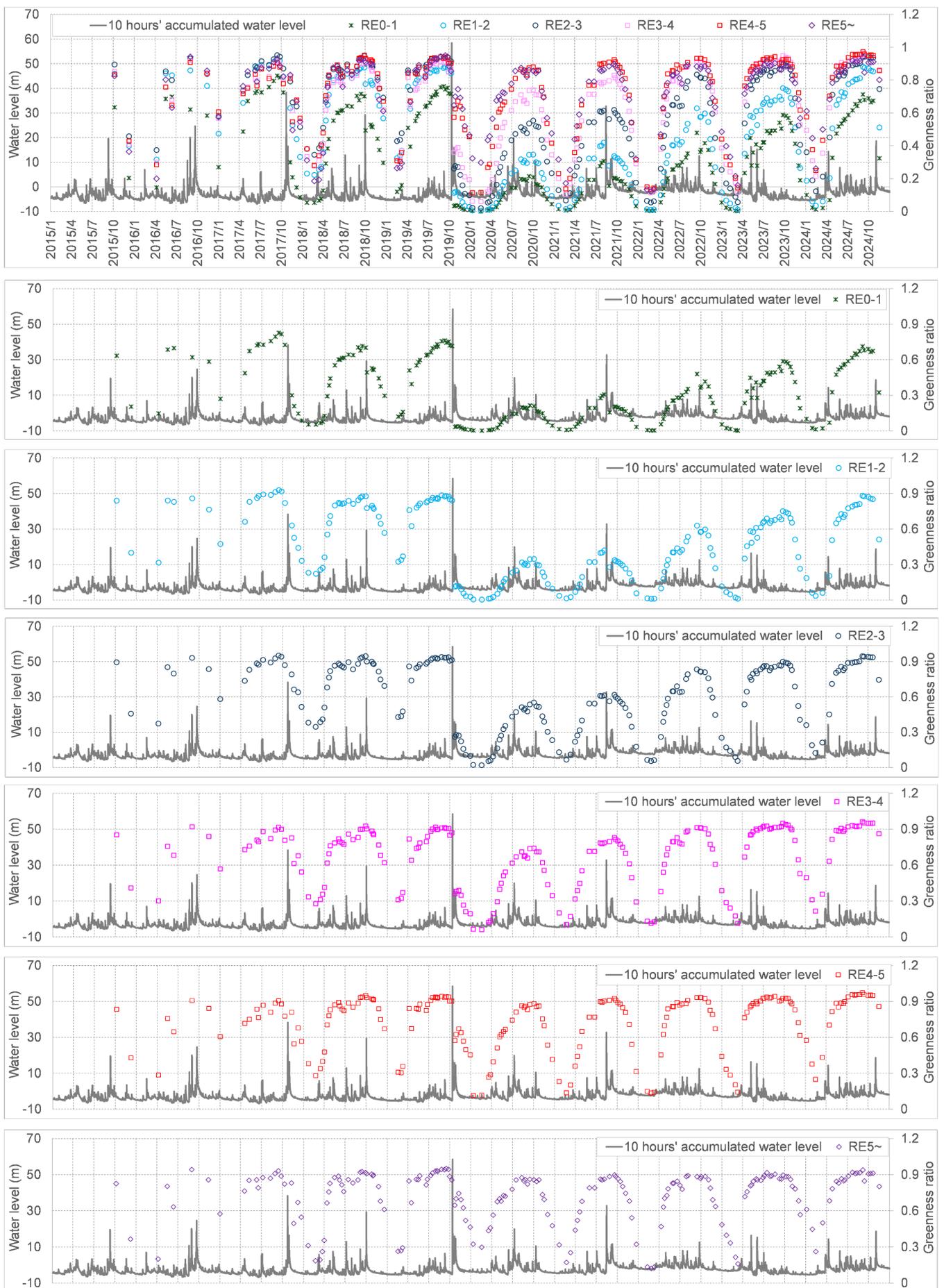


Figure 8: Spatiotemporal vegetation greenness and hydro-morphological dynamics (RE0-1, RE1-2, RE2-3, RE3-4, RE4-5, and RE5~ denote relative elevation classes of 0–1 m, 1–2 m, 2–3 m, 3–4 m, 4–5 m, and ≥ 5 m, respectively. “10-hour accumulated water level” refers to the water level accumulated over a 10-hour period).

230 **3.2 The annual maximum greenness ratio and NDVI along the channel-lateral transect**

We further examined the temporal dynamics of riparian vegetation greenness over a 9-year period from 2016 to 2024, excluding 2015 due to limited NDVI data availability. The analysis focused on annual maximum NDVI across varying relative elevations. The maximum greenness ratio was extracted and analyzed because NDVI typically peaks between July and October at the study site, the primary growing season for both herbaceous and woody vegetation (Pettorelli et al., 2005). During this
 235 period, the annual maximum greenness ratio was used as a proxy for peak vegetation cover, because vegetation was fully developed (NDVI > 0.2).

Figure 9 illustrates the lateral distribution of the greenness ratio, measured from the shoreline (relative elevation = 0 m) to the levee bank (above 5 m). A clear increasing trend in the greenness ratio was observed from the river’s edge up to approximately 5 m in relative elevation. The relationship between greenness ratio and relative elevation was modelled using an empirical function (Eq. (4)), with the fitted curves for each year shown in **Figure 9**. The parameters L , h_0 and k utilized in the model
 240 are summarized in **Table 1**. Among the parameters, h_0 , which indicates the relative elevation of the greenness ratio to be reached at 50%, was found to vary with flood conditions from the preceding year.

$$Green(h) = L / (1 + \exp(-k(h - h_0))) \tag{4}$$

$G(h)$: Greenness ratio with a relationship to relative elevation (h)

245 L : Maximum greenness ratio

h_0 : the relative elevation at which the greenness ratio equals 0.5 (50%)

k : Steepness parameter

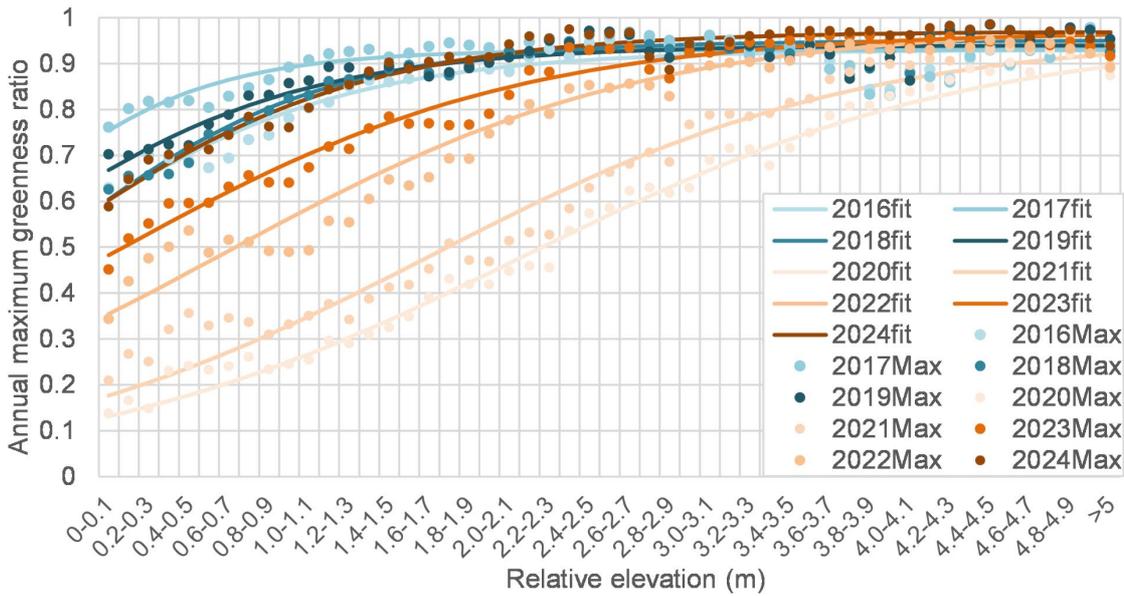


Figure 9: Annual maximum greenness ratio along the transect.

Table 1: Parameters utilized in the fitting curve.

Parameter	2016	2017	2018	2019	2020	2021	2022	2023	2024
L	0.93	0.93	0.95	0.94	0.96	0.96	0.98	0.97	0.97
h_0	-0.42	-0.72	-0.34	-0.63	2.12	1.65	0.64	0.06	-0.35
k	1.3	1.89	1.43	1.32	0.89	0.93	0.96	0.98	1.24

Figure 10 shows annual maximum NDVI along the transect at different relative elevations. Distinct differences in NDVI distribution patterns along transects were observed between the periods 2016–2019 and 2020–2024. During 2016–2019, NDVI
 255 values increased with relative elevation from 0 m up to approximately 3 m, and then decreased between 3 and 5 m. Prominent

NDVI peaks were identified around the zone between 1.5 and 3.0 m. In contrast, during 2020–2024, NDVI exhibited a more gradual and consistent increase from 0 to 5 m, indicating a different spatial pattern of vegetation greenness ratio.

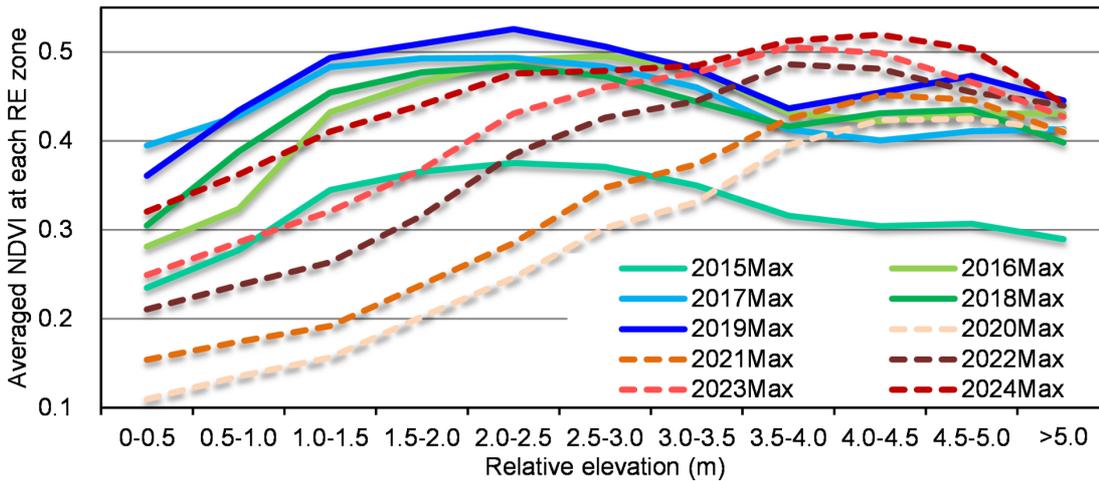


Figure 10: Annual maximum NDVI averaged within each 0.5m relative elevation zone along the transect.

260 3.3 Seasonal dynamics of riparian greenness

We computed the monthly mean greenness ratio for the period January 2016–November 2024, using NDVI values extracted from approximately 200 Sentinel-2 images. Since only two images were available for 2015, that year was excluded from the analysis. From March to June, the greenness ratio steadily increased with rising temperatures, and precipitation created favorable conditions for photosynthesis. However, the upward trend slowed during July and August, a trend that may have been caused by elevated temperatures and high cumulative precipitation that stimulated respiratory activity more than photosynthesis (Fig. 11). As temperatures and rainfall moderated from mid-summer peaks, riparian vegetation entered a secondary growth phase, and the greenness index reached a maximum during September (Wang et al., 2025). A post-October decline in the greenness index correlated to reduced temperatures and precipitation regimes, leading to diminished photosynthetic activity.

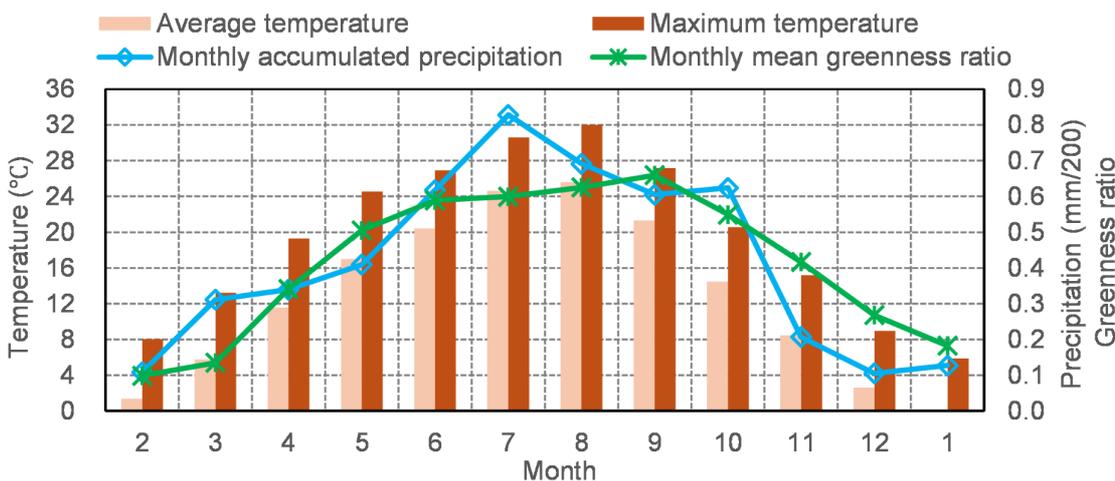


Figure 11: The relationship between seasonal dynamics of greenness ratio and temperature and precipitation.

4 Discussions

4.1 Spatiotemporal Sentinel-2 NDVI response to flood-disturbance and lateral elevation change

4.1.1 Sentinel-2 NDVI response to varying flood magnitudes

275 Our results provide quantitative evidence supporting the biogeomorphic succession framework by demonstrating how
 riparian vegetation responses diverge across flood pulses of different magnitudes. **Figure 13** illustrates the annual maximum
 greenness ratio and the parameter h_0 (the relative elevation where the greenness ratio equals 0.5) against annual maximum
 water levels. Notably, sharp declines in maximum NDVI during 2018 and 2020 reflect substantial vegetation removal caused
 by the high-magnitude floods of 2017 and 2019, respectively. A sharp decline in h_0 was also observed during 2020,
 280 indicating that riparian vegetation was likely removed at a lower relative elevation, correspondingly with the effect of the
 mega flood in 2019. High-magnitude events (specifically those with return periods of more than five years) tend to cause
 widespread scouring (Nallaperuma and Asaeda, 2020) and drastic channel reshaping (**Fig. 4**), effectively resetting the
 riparian distribution and structure (Death et al., 2015; Stella et al., 2013). The post-2019 increase in the logistic parameter h_0 ,
 which shifted from negative values to over 2.0 m, quantifies a significant recovery of vegetation in lower elevation zones.
 285 This "clearing the template" effect aligns with previous research (Kyuka et al., 2020), where extreme external forces during
 mega-floods overcome root stabilization and resets the successional cycle. Conversely, at higher elevations (3.5–4.5 m),
 NDVI values exceeded pre-flood levels (**Fig.10**). This suggests that the mega-flood provided a vital subsidy of nutrients and
 moisture to high-elevation zones typically disconnected during ordinary flow events (Džubáková et al., 2015), thereby
 stimulating productivity in previously water-limited areas.

290 In contrast, while vegetation reduction occurred between 0–3 m during the 2018 and 2021 events (**Fig. 8**), the annual maximum
 greenness ratio increased and h_0 decreased (**Fig. 12**) in the subsequent years (2019 and 2022). This pattern suggests rapid post-
 flood regrowth following moderate floods. In these scenarios, pioneer species quickly exploit nutrient-rich sediment and open
 space created by the flood pulse (Camporeale et al., 2006; Greet et al., 2011), illustrating a more resilient and localized
 successional response compared to the transformative impacts of mega-floods.

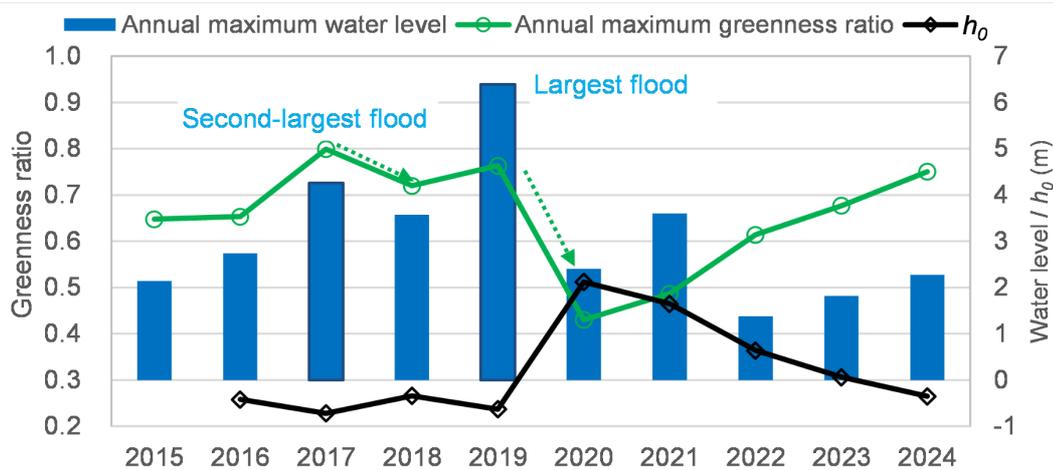


Figure 12: Relationship between annual maximum greenness / h_0 and annual maximum water level (h_0 : the relative elevation at which the greenness ratio equals 0.5).

4.1.2 Riparian vegetation dynamics along the lateral transect zone

Based on the 10-year dynamics of greenness ratio and NDVI (**Fig.8**, **Fig.9** and **Fig.10**), we identify three functional zones of
 300 riparian vegetation by their response to disturbance and resource availability:

- (1) The high-disturbance zone (relative elevation (RE) ≤ 1.0 m): Vegetation in this zone exists in a "perpetual pioneer" state. Frequent small-to-moderate floods (return periods ≤ 5 years) prevent woody encroachment, supporting the intermediate disturbance hypothesis (IDH) (Connell 1978). Here, the frequency of physical disturbance is too high for vegetation to reach the "biogeomorphic succession" where it could significantly alter channel morphodynamics (Corenblit et al., 2007).
 305 Consequently, this zone acts as a transient environment where biomass accumulation is strictly limited by frequent disturbance.

(2) The biogeomorphic feedback zone ($1.0\text{m} < \text{RE} \leq 3.5\text{m}$): This band represents a critical "window of opportunity" (Balke et al., 2014) for woody species like *Robinia pseudoacacia* to survive the seedling stage. This zone acts as an optimal ecogeomorphic niche, where moderate disturbance and nutrient subsidies supported the high annual NDVI peaks observed before 2019 (Fig. 10). Upon transitioning to the biogeomorphic phase (Corenblit et al., 2007), these species act as ecosystem engineers (Jones et al., 1994) by trapping sediment and stabilizing landforms. While this promotes aggradation, it simultaneously elevates flood risk by increasing hydraulic roughness and reducing the channel's conveyance capacity (Oorschot et al., 2016).

(3) The stable zone ($\text{RE} > 3.5\text{m}$): At these higher elevations, vegetation is disconnected from regular hydromorphological changes. While these plants exhibit high resistance to most floods, they may experience moisture stress due to insufficient inundation, which can constrain growth rates compared to the intermediate zone (Marchetti et al., 2020; Modi et al., 2022). Reduced water and nutrient availability are likely to contribute to the lower NDVI values observed in the stable zone (Fig.10). Following the 2019 mega-flood, particularly during 2021–2024, the elevation of annual maximum NDVI shifted from the 1.5–3.0 m band to the 3.5–4.5 m band (Fig.10), indicating a potential upward migration of the biogeomorphic feedback zone. This shift suggests that vegetation–morphology interactions may have reorganized toward higher elevations after the extreme disturbance (Balke et al., 2014). However, the absence of major flood events since 2021 limits our ability to evaluate how subsequent disturbance regimes will influence the persistence and stability of these newly established vegetation patterns. Previous studies have emphasized that stochastic external forcing, such as infrequent extreme floods, plays a critical role in driving ecosystem transitions. Accordingly, longer-term observations are required to determine whether the 2019 mega-flood permanently redefined the boundaries among high-disturbance, biogeomorphic feedback, and stable zones, or whether the observed elevation shift represents a transient successional response rather than a regime shift.

4.2 Practicality of Sentinel-2 NDVI for frequent-disturbed riparian environment

The application of Sentinel-2 NDVI monitoring in this study demonstrates high practicality for adaptive river management. The primary advantage lies in its high temporal resolution, which allows for capturing the "resetting" of vegetation immediately after flood pulses and the subsequent recovery processes, dynamics that are typically missed by traditional annual or biennial aerial surveys (Gurnell et al., 2003; Bertoldi et al., 2011). As shown in Figure 8 and Figure 12, NDVI and greenness ratio responses are jointly controlled by flood magnitude (linked to return period) and relative elevation, with slight and pronounced declines following floods with approximately 10-year and 50-year return periods, respectively. Beyond disturbance detection, post-flood NDVI recovery patterns provide quantitative insight into vegetation recruitment and growth rates. For example, the faster recovery observed between 2021 and 2022 compared to 2020–2021 may offer empirical constraints for calibrating process-based riparian vegetation models (Nallaperuma and Asaeda, 2020; Toda et al., 2020). The above results confirm that Sentinel-2–derived NDVI is a robust indicator for capturing the dual nature of riparian dynamics (abrupt flood-induced removal and gradual successional recovery) and their responses to hydromorphological change. It possesses high potential for application in more stable or other high-disturbance riparian environments. Nevertheless, the specific physical mechanisms of vegetation destruction (e.g., uprooting, burial, bending, or shear stress thresholds) cannot be analyzed directly within the current NDVI-based framework and require complementary field or process-based analyses.

Except the potential of Sentinel-2 NDVI on presenting the relationship between vegetation dynamics and hydrodynamics. This database may also have the potential for the analysis of vegetation community distribution. The greenness ratio reflects the presence or absence of vegetation based on an NDVI threshold (> 0.2), serving as a binary indicator of vegetated area. In contrast, NDVI values themselves provide more detailed information on vegetation health and type, with higher values typically indicating denser and healthier vegetation. At peak summer greenness, trees generally show higher NDVI than grasses due to greater leaf area and NIR reflectance; accordingly, NDVI helps discriminate woody versus herbaceous vegetation here. Prominent NDVI peaks were identified around the zone between 1.5 and 3.0 m (Fig.10), however, the clear peak of greenness ratio was not identified in Figure 9. In this study, the ratio of each vegetation type along the transect was

calculated based on Environment Information Map 2018 and relative elevation. Environment Information Map 2018 records the vegetation species distribution and was surveyed by the MLIT. As shown in **Figure 13**, grass-dominated zones (green line) and tree-dominated zones (black line) spatially differed during 2018, which means that the environmental optima of the different vegetation species are different along the lateral environmental gradients (Camporeale and Ridolfi, 2006; Johnson et al., 1995). At the current study site, *Robinia pseudoacacia*, the dominant riparian tree species, is primarily distributed between 2.0 and 4.0 m in relative elevation. Since riparian trees typically exhibit higher NDVI values than grasses, the concentration of *Robinia pseudoacacia* in this elevation range likely contributed to the distinct NDVI peaks observed from 2016–2019. In the previous research, the Sentinel-2 time-series imagery was reported to have the possibility on grassland plant communities mapping (Rapinel et al., 2019), it may also have the potential to be utilized as the vegetation community’s analysis for the frequent disturbed environments based on the spatial greenness ratio and NDVI distribution.

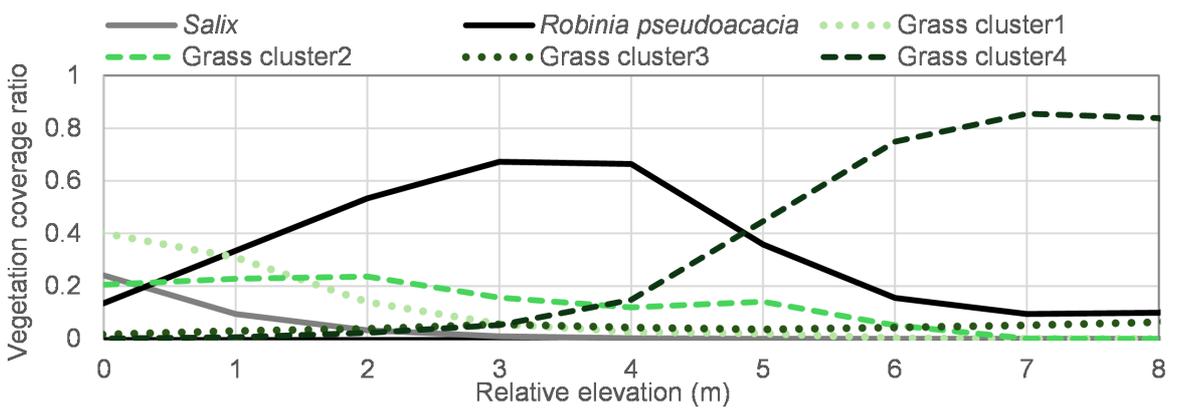


Figure 13: Vegetation type and its coverage ratio along the transect.

4.3 Management implications: optimizing the "where" and "when"

Effective riparian vegetation management requires addressing two key dimensions: identifying the optimal timing (“when”) and the spatial prioritization (“where”) of interventions. By integrating our NDVI-derived insights with biogeomorphic theory, we propose a strategic framework that balances flood conveyance needs with ecosystem stability.

4.3.1 The timing of vegetation management

Our results indicate that high-magnitude floods (e.g., the 2017 and 2019 events) act as a natural “resetting” mechanism by causing widespread vegetation removal. In contrast, smaller and moderate floods (return periods ≤ 5 years) produce limited disturbance and are often followed by rapid vegetation expansion in the subsequent growing season, as evidenced by the sharp increase in greenness ratio observed from 2021 to 2022. Although the greenness ratio declined substantially after the 2019 event, the steady recovery observed during 2020–2024, when greenness ratio nearly returned to pre-flood levels, highlights the high resilience of the riparian vegetation system. From a management perspective, higher greenness ratio and NDVI values, reflecting greater vegetation cover and denser biomass, are associated with increased hydraulic roughness. Increased roughness reduces flow conveyance and may elevate the risk of embankment overtopping during floods (Al Mehedi et al., 2024). Therefore, vegetation management should be considered not only in relation to small-to-moderate floods (return periods ≤ 5 years), but also in the context of post-extreme-flood recovery, and continuous post-flood monitoring is essential to track regrowth trajectories and support adaptive intervention strategies.

At the Chikuma River site, NDVI values typically peak in September, coinciding with the end of the growing season. However, major flood events in this region, often triggered by typhoons, tend to occur in late September or October. This seasonal overlap implies that peak vegetation development aligns with the onset of extreme flood events, posing elevated risks to

embankment overflow. Therefore, proactive management, such as selective cutting, should be prioritized in August, prior to the onset of the typhoon season.

4.3.2 Spatial prioritization for management actions

The lateral variation of NDVI and the greenness ratio (Fig. 14) reveals how management should be prioritized across different elevational zones: Low-elevation zones: While vegetation here recovers rapidly post-disturbance, it is frequently washed out by moderate floods. Management in this zone is largely "self-maintained" by natural fluvial processes (Death et al., 2015), requiring minimal human intervention except for removing large debris. High-elevation zones: These areas harbor stable vegetation that resists most flood events. While this vegetation provides bank stability, it can significantly obstruct flow during infrequent, high-magnitude floods. Management here should involve periodic, long-term strategic thinning to ensure conveyance during extreme events (Herrmann et al., 2024). Intermediate-elevation zones: This zone represents the most critical area for targeted management. As observed in our NDVI profiles (Fig.10), these elevations support moderate-to-high density vegetation that impacts flow capacity for both medium and large floods. This aligns with findings by Marchetti et al. (2020) and Modi et al. (2022), suggesting that focusing interventions on these intermediate heights provides the most balanced reduction in flood risk. By utilizing the h_0 parameter (relative elevation at which the vegetation coverage equals 0.5) as a monitoring threshold, river managers may quantitatively identify when the "vegetation expansion front" has advanced too far into the channel, triggering evidence-based, spatially targeted cutting operations that maximize efficiency on flood management.

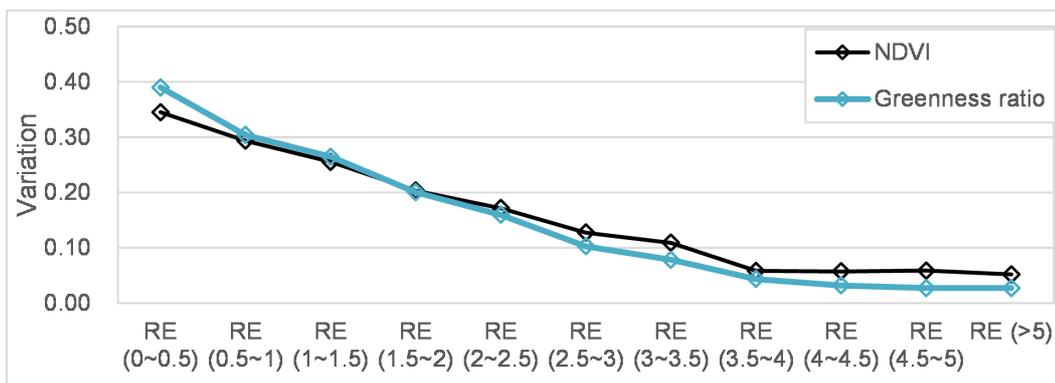


Figure 14: Averaged variation of NDVI and greenness ratio from 2016 to 2024.

From the above analysis, spatiotemporal NDVI dynamics and their responses to hydrological disturbances provide a quantitative basis for identifying the timing and necessity of vegetation interventions. The response of riparian vegetation on different disturbances and relative elevation zones was consistent with the intermediate disturbance hypothesis and lateral conceptual theory proposed in previous studies. From a river-management perspective, efficient and cost-effective strategies are essential. Vegetation cutting in low relative-elevation zones may be economically inefficient because vegetation in these areas is frequently removed by natural flood disturbances. In contrast, interventions in stable, high-elevation zones may also have limited efficiency, as vegetation there is affected primarily during extreme floods and contributes less to conveyance under more frequent events. The framework developed in this study (integrating spatiotemporal NDVI dynamics with flood disturbances and relative elevation) therefore provides a practical basis for prioritizing where and when vegetation management should be implemented to maximize benefits while minimizing unnecessary costs. However, this database may not be applied to river systems with narrower channels with the limitation of their spatial resolution.

410 **5 Conclusions**

This study utilized a decade of Sentinel-2 imagery and NDVI to evaluate riparian vegetation dynamics in the Chikuma River, bridging the gap between remote sensing and river management under frequent disturbances. The key conclusions are as follows:

415 First, the response of NDVI is highly sensitive to flood magnitude. While moderate floods caused only slight NDVI declines, high-magnitude events like the 2019 mega-flood acted as a primary resetting mechanism. Along the lateral elevation gradient, three functional zones (high-disturbance, biogeomorphic feedback, and stable) were identified, with the intermediate band (~1.5–3.5 m) showing the high potential for woody vegetation succession.

Second, Dense Sentinel-2 time series captured both abrupt post-flood vegetation removal and multi-year recovery trajectories that are often missed by annual and biennial surveys.

420 Finally, we propose a practical framework to optimize management "when" and "where". Timing ("when") should prioritize post-event actions after intermediate floods (return period ≤ 5 years). These floods may not fully reset vegetation, but rapid regrowth can quickly increase hydraulic roughness. Seasonal actions should also be implemented in August to maximize roughness reduction before the typhoon season.

Author contribution

425 YZ planned the study and methodology; YZ and RZ performed the investigation; YZ conducted the formal analysis and curated the data (with additional data curation by YT); YZ wrote the manuscript draft; YT and RZ reviewed and edited the manuscript; YT provided validation and acquired funding.

Competing interests

430 The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

Data availability

435 Sentinel-2 imagery used in this study is available at <https://doi.org/10.5281/zenodo.17240522> (ZHOU, 2025). The temporal water level was download from the <https://www1.river.go.jp/>. The temperature and precipitation data were download from <https://www.data.jma.go.jp/stats/etrn/>. The vegetation type data was downloaded from the <https://www.nilim.go.jp/lab/fbg/ksnkankyo/>. The temporal water level, temperature and precipitation data are available at <https://doi.org/10.5281/zenodo.18092794> (ZHOU, 2025). Topography data of ALB and LP will be made available on request.

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