

# 1 **Modification and validation of a commercial dynamic chamber for reactive** 2 **nitrogen and greenhouse gas flux measurements**

3 Moxy Shah<sup>1</sup>, Kifle Z. Aregahegn<sup>1</sup>, Danial Nodeh-Farahani<sup>1</sup>, Leigh R. Crilley<sup>1,‡</sup>, Tasnia Hasan<sup>1</sup>,  
4 Yashar Ebrahimi-Iranpour<sup>1</sup>, Fahim Sarker<sup>1</sup>, Nick Nickerson<sup>2</sup>, Chance Creelman<sup>2</sup>, Sarah Ellis<sup>2</sup>,  
5 Alexander Moravek<sup>1,§</sup>, Trevor C. VandenBoer<sup>1,\*</sup>

6  
7 <sup>1</sup> Department of Chemistry, York University, Toronto, Ontario, Canada

8 <sup>2</sup> Eosense Inc., Dartmouth, Nova Scotia, Canada

9 <sup>‡</sup> Now at: Atmospheric Services, WSP Australia, Brisbane, QLD, Australia

10 <sup>§</sup> Now at: German Environment Agency, Department of Air Quality, Dessau-Rosslau, Germany

11  
12 \*Communicating author: tvandenb@yorku.ca

## 13 14 15 **Abstract**

16 Reactive nitrogen compounds (NO, NO<sub>2</sub>, HONO, NH<sub>3</sub> and others; N<sub>r</sub>) play important roles in  
17 atmospheric processes, and their cascading impacts throughout the Earth system have adverse  
18 effects on both the environment and human health. The fluxes of these gases at the surface-  
19 atmosphere interface have been studied in isolation or ~~in smaller subsets~~ ~~by micrometeorological~~  
20 ~~techniques or chambers~~, but simultaneous ~~observations~~ ~~fluxes~~ of all N<sub>r</sub> ~~species~~ alongside standard  
21 greenhouse gases (GHGs) ~~as a function of time~~ have not been reported. Here, a dual-dynamic  
22 chamber system was developed for N<sub>r</sub> by modifying a commercially available system for GHG  
23 fluxes for use with destructive analyzers, ~~and including a reference chamber~~ to account for  
24 chemical ~~changes~~ ~~challenges~~. The resulting platform makes the measurement of N<sub>r</sub> and, by  
25 extension other reactive gases, more widely accessible to the scientific community, as custom  
26 chambers do not need to be fabricated.

27  
28 System modifications to passivate surfaces ~~were implemented, so that N<sub>r</sub> gases like NO<sub>2</sub> could be~~  
29 ~~effectively transferred to standard gas analyzers, with~~ ~~reduced~~ an initial 36% loss ~~of NO<sub>2</sub>~~ due to  
30 transformations ~~ultimately minimized~~ below analyzer detection limits (~10%) ~~under for~~ relevant  
31 atmospheric conditions. The modified 72 L chamber ~~response times~~ did not ~~see a change in the~~

32 ~~baseline response times~~ for GHGs or NO at a flow rate of 2 L min<sup>-1</sup>. ~~They retained the same values~~  
33 ~~as an ideal non-reactive trace gas~~ ( $\tau = 37\text{-}39$  min versus 36 min). The modifications improved the  
34 transfer ~~time constants~~ of NO<sub>2</sub>, HONO, and NH<sub>3</sub> by up to 2 min, but substantial surface interactions  
35 for NH<sub>3</sub> remain. ~~In all cases, a~~ surface interaction term ~~needs to be~~ was characterized for these  
36 gases to obtain accurate field fluxes. ~~Losses of NO<sub>2</sub> and O<sub>3</sub> by known gas phase reactions, or from~~  
37 ~~deposition and reaction on pristine and aged chamber surfaces, were characterized across a range~~  
38 ~~of environmentally relevant relative humidities (RH) and mixing ratios. The final dual-chamber~~  
39 ~~system configuration includes a measurement and reference chamber, which are necessary to~~  
40 ~~implement the corrections for surface effects and chemical transformations when accurately~~  
41 ~~quantifying dynamic fluxes~~ via a mass balance framework.

42  
43 Proof-of-concept measurements of N<sub>r</sub> fluxes from agricultural soil samples under controlled lab  
44 conditions as a function of soil water content were able to quantify emissions of NO, NO<sub>2</sub>, HONO,  
45 NH<sub>3</sub>, and N<sub>2</sub>O simultaneously, without amendment and when subject to N<sub>r</sub> fertilization  
46 ~~experiments using urea, ammonium carbonate and bicarbonate, and ammonium nitrate.~~  
47 Unfertilized ~~replicate agricultural soils~~ samples showed variability in NO<sub>2</sub> and HONO emissions  
48 when ~~prepared with minimal disturbance to the~~ soil structure was minimally disturbed, consistent  
49 with ~~values consistent with those reported by~~ in-situ field measurements from others. These oppose  
50 maximum potential fluxes ~~characterized~~ in prior lab soil manipulations, particularly for HONO  
51 relative to NO. Last, N<sub>r</sub> field fluxes were quantified ~~with destructive gas analyzers in the field~~ with  
52 the dual-chamber system on an in-use agricultural soil, and including a urea-based fertilizer  
53 perturbation to stimulate microbial and chemical transformation and transfer N<sub>r</sub> to the atmosphere.  
54 ~~The resulting fluxes observed show g~~ Good agreement with ~~prior reports based on~~ other field flux  
55 techniques was found. The mass balance terms within the dual-chamber approach are fully  
56 inspected from the pilot deployment in the field, along with an error analysis, to aid in the uptake  
57 of this approach by the community.

58

## 59 1 Introduction

60 ~~The Earth's biogeochemical~~ nitrogen (N) cycle is ~~a crucial biogeochemical cycle on Earth,~~  
61 essential for sustaining life through the production of nucleic acids, proteins, and other vital  
62 biomolecules (Lehnert et al., 2021). ~~Although t~~The carbon cycle receives ~~the most attention~~much  
63 focus due to ~~looming the~~ ~~climate~~ ~~crises caused by~~ ~~impacts~~ greenhouse gases (GHGs) like carbon  
64 dioxide (CO<sub>2</sub>) and methane (CH<sub>4</sub>), ~~yet~~ the N cycle is ~~closely~~ intertwined ~~with the carbon cycle~~  
65 (Schlesinger, 2020). At the interface of these cycles and the Earth's surface, reactive nitrogen (N<sub>r</sub>)  
66 species exchanged between ecosystems and the atmosphere have therefore become an area of  
67 emerging interest (Lehnert et al., 2021; Wu et al., 2020). ~~Within the atmosphere,~~ Atmospheric N<sub>r</sub>  
68 species such as nitric oxide (NO) and nitrogen dioxide (NO<sub>2</sub>) – collectively referred to as NO<sub>x</sub> –  
69 ammonia (NH<sub>3</sub>), and nitrous acid (HONO) can ~~each be mediated in part through~~ ~~experience~~  
70 surface-atmosphere exchange, ~~influencing~~ impacting local air or water quality, ecosystem  
71 processes, and biodiversity (Lehnert et al., 2021; Richardson et al., 2023; Wu et al., 2020).  
72 Meanwhile, the non-reactive nitrous oxide (N<sub>2</sub>O) has climate impacts due to its long ~120 year  
73 atmospheric lifetime (~~~120 years~~) and ~~potent greenhouse effect~~ (IPCC, 2023).

74 Reactive nitrogen ~~compounds-gases~~ (NO, NO<sub>2</sub>, HONO, and NH<sub>3</sub>; N<sub>r</sub>) play ~~an~~ important roles in  
75 atmospheric processes, contributing to the formation of pollutants like ozone (O<sub>3</sub>) and secondary  
76 organic aerosols (SOA). The exchange of N<sub>r</sub> between the Earth's surface and the atmosphere  
77 involves ~~numerous~~ production and loss processes driven by both natural and human activities.  
78 ~~Reactive nitrogen species~~ They are removed ~~from the atmosphere~~ through wet and dry deposition,  
79 and their abundance ~~depends on~~ reflects the net outcome relative to N<sub>r</sub> emissions (Delaria & Cohen,  
80 2023). At the land surface, N<sub>r</sub> is released into the atmosphere by microbial nitrogen cycling,  
81 agricultural activities, wildfires, or fossil fuel combustion (Benedict et al., 2017; Mosier, 2008;  
82 Yang et al., 2024). ~~Despite the wealth of knowledge on the environmental importance of N<sub>r</sub>,~~  
83 sStudying N<sub>r</sub> at the surface-atmosphere interface with high time resolution and chemical speciation  
84 remains a challenge due to its high spatial and temporal variability driven by factors like climate,  
85 vegetation cover, and soil/surface properties (Ludwig et al., 2001). For example, ~~recent studies~~  
86 have demonstrated that vertically resolved HONO production observed at the ground surface  
87 demonstrated how it plays a major role in the unexplained daytime HONO source and its impact  
88 on daytime OH, ~~with vertically resolved measurements confirming that the surface remains a~~  
89 dominant source at all times of day (VandenBoer et al., 2013, 2015; Young et al., 2012). Such

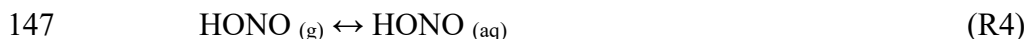
90 observations pose a challenge ~~in studying N<sub>r</sub>~~ because suitable ~~equipment for~~ high time resolution  
91 ~~atmospheric observations equipment~~ is expensive, ~~limiting the concurrent study of preventing~~ the  
92 interplay between emission and deposition for ~~many all~~ N<sub>r</sub> species ~~to one or a few species at a~~  
93 ~~time from being concurrently studied~~. As a result, no systems are ~~currently sufficiently~~ accessible  
94 ~~enough~~ to the scientific community ~~that they may to~~ be deployed ~~aeross~~-widely ~~across~~ different  
95 global landscapes, ~~but~~ particularly soils.

96 Soils ~~play a pivotal role in mediating N<sub>r</sub> emissions due to their~~have a dual function as both a source  
97 and a sink of ~~nitrogenous N<sub>r</sub>~~ species. Soil-atmosphere exchange of N<sub>r</sub> is thought to be governed  
98 by ~~atmospheric abundance and/or soil~~ microbial processes such as nitrification and denitrification,  
99 with ~~soil~~-factors like pH, moisture, organic matter, and nitrogen availability ~~playing a crucial role~~  
100 ~~in~~ regulating ~~N<sub>r</sub> emission flux directionality~~ (Mosier, 2008; Purchase et al., 2023; Stepniewski et  
101 al., 2015). ~~These m~~Microbial processes have been demonstrated to drive the formation and release  
102 of ~~some~~ N<sub>r</sub> species ~~(e.g., like~~ NO, NH<sub>3</sub> and N<sub>2</sub>O) ~~from soils, in addition to and have~~ -assertions ~~that~~  
103 ~~they can support the production and release of with respect to~~ HONO ~~via partitioning~~ (Butterbach-  
104 Bahl & Dannenmann, 2011; Kool et al., 2010; Mushinski et al., 2019; Oswald et al., 2013; Su et  
105 al., 2011). Understanding the exchange of ~~these all~~ N<sub>r</sub> gases is essential for unravelling the complex  
106 interactions between the nitrogen and carbon cycles and their broader environmental impacts from  
107 an unprecedented imbalance in the global nitrogen cycle (Richardson et al., 2023). ~~Given the~~  
108 ~~importance of N<sub>r</sub> in atmospheric chemistry and its influence on ecosystem processes, it is essential~~  
109 ~~to understand the mechanisms driving its exchange~~ (Fowler et al., 2013). ~~To date, comprehensive~~  
110 ~~exchange systems for studying the full suite of N<sub>r</sub> have not been reported.~~

111  
112 To quantify N<sub>r</sub> exchange for a few explicit species near the ground surface, various flux  
113 measurement techniques have been employed, ~~yielding insights into their magnitude and~~  
114 ~~identifying some driving factors.~~ Quantifying fluxes of atmospheric gases has been most  
115 effectively applied to GHGs at the surface-atmosphere interface. Traditional measurement  
116 methods, such as eddy covariance (EC), relaxed eddy accumulation (REA), and aerodynamic  
117 gradient (AG) methods, have been extensively used for ecosystem-scale, continuous flux  
118 monitoring, including targeted assessment of most N<sub>r</sub> species (e.g., ~~(~~Bao et al., 2022; Geddes &  
119 Murphy, 2014; Kamp et al., 2020; Laufs et al., 2017; Min et al., 2014; Moravek et al., 2014, 2019;

120 Ren et al., 2011; Von Der Heyden et al., 2022; Wang et al., 2022; Wolff et al., 2010). These  
121 micrometeorological techniques measure concentrations, concentration gradients, and/or  
122 turbulence to estimate fluxes across interfaces applicable to ecosystem-scale processes. When  
123 operated continuously, they offer long-term insight without disrupting the natural system.  
124 Chamber methods have some advantages compared to micrometeorological techniques, as they  
125 are relatively inexpensive, easy to deploy, and require minimal prior meteorological training and  
126 expertise (Tang et al., 2020). Chambers are limited to small plots, e.g. making them suitable to  
127 study the effect of different fertilizer treatments (Anthony & Silver, 2024; Chiaravalloti et al.,  
128 2023; Manco et al., 2025; Tang et al., 2020). The chamber method is a widely used technique for  
129 measuring GHG fluxes, especially for N<sub>2</sub>O. However, challenges remain in addressing potential  
130 biases introduced by chemical transformations occurring within the chamber and interactions with  
131 the chamber surfaces, particularly for N<sub>r</sub> species (i.e., NO<sub>2</sub>, HONO, NH<sub>3</sub>).

132 Chemical transformations can occur on chamber surfaces or between the point of emission and  
133 measurement, where ~~S~~ surface interactions such as adsorption, desorption, and heterogeneous  
134 reactions can alter the apparent concentration of several N<sub>r</sub> species. For example, NO in a chamber  
135 could react ~~in the gas phase~~ with O<sub>3</sub> to form NO<sub>2</sub> during the day and NO<sub>3</sub> at night if sufficient O<sub>3</sub>  
136 is present to fully titrate NO (R1, R2). Heterogeneous reaction of gas-phase NO<sub>2</sub> on surfaces under  
137 humid conditions also produces nitric acid (HNO<sub>3</sub>) and HONO (R3) (Kleffmann et al., 2005;  
138 Ramazan et al., 2004; VandenBoer et al., 2015). The formed HONO can undergo multiphase  
139 processes in the chamber or on enclosed surfaces by partitioning into water according to its  
140 Henry's Law constant and then dissociating into nitrite (NO<sub>2</sub><sup>-</sup>) and the hydronium ion (H<sub>3</sub>O<sup>+</sup>)  
141 according to its acid dissociation equilibrium constant and the pH (R4, R5) (He et al., 2006; Ren  
142 et al., 2020). Nitrous acid could also photolyze, yielding NO and an OH radical (R6) (Spataro &  
143 Ianniello, 2014).





150 These processes in/on the chamber can introduce uncertainty in flux measurements. Characterizing  
151 and accounting for chemistry and surface effects in chamber-based flux methods are therefore  
152 ~~necessary to obtain accurate flux measurements of  $\text{N}_r$  species with a chamber-based system. While~~  
153 sStatic chamber systems typically determine the flux from the change in headspace concentration  
154 after closing the ~~chamber cover~~, ~~traditional d~~Dynamic chamber systems use have traditionally  
155 used a controlled flow of ambient air through the headspace and to retrieve the flux from ~~the a~~  
156 concentration difference at between the chamber inlet and outlet. ~~With minor modifications, these~~  
157 ~~chambers have previously been applied to agricultural soils and natural ecosystems, demonstrating~~  
158 ~~their utility for measuring emissions of gases such as  $\text{N}_2\text{O}$  and  $\text{NO}_x$  (Maggiotto et al., 2000; Pape~~  
159 ~~et al., 2009). Their ability to create standardized and reproducible conditions makes them~~  
160 ~~particularly advantageous for studying the driving mechanisms behind gas exchange, yet their~~  
161 ~~application to the full suite of relevant  $\text{N}_r$  species for surface-atmosphere exchange remains~~  
162 ~~unestablished.~~

163 The dynamic chamber flux method has ~~been reported for measurements of the exchange fluxes of~~  
164 ~~measured~~ challenging gases like biogenic volatile organic compounds ~~(BVOCs, such as~~  
165 ~~monoterpenes and isoprene,) from vegetation and farmland (Kolari et al., 2012; Mochizuki et al.,~~  
166 ~~2018), Pugliese et al. 2023),  $\text{NH}_3$  volatilization from cattle manure (Becciolini et al., 2024),  $\text{N}_2\text{O}$~~   
167 ~~and  $\text{NO}_x$  from turfgrass (Maggiotto et al., 2000), and  $\text{NO}_x$  fluxes from grasslands (Pape et al.,~~  
168 ~~2009; Plake et al., 2015). Scharko et al. (2015) used sealed chambers, while and Tang et al. (2019),~~  
169 ~~using used~~ dynamic ~~chambers, were amongst the first~~ to highlight the hotspot potential for both  
170  $\text{HONO}$  and  $\text{NO}_x$  fluxes from agricultural soils ~~as hotspots impacting atmospheric chemistry, which~~  
171 ~~has been implemented in atmospheric models (Ha et al., 2023; Tian et al., 2024). These are of~~  
172 high interest due to their impacts on atmospheric chemistry from local to regional scales.

173 The complexity biological and chemical controls of n nitrite ( $\text{NO}_2^-$ ) ~~chemical and biological~~  
174 production and loss in soils, coupled with soil properties facilitating gas exchange of  $\text{HONO}$ , has  
175 led to intense interest and debate around discerning the fundamental controls on its surface-  
176 atmosphere exchange (R4, R5) (Barney & Finlayson-Pitts, 2000; Huang et al., 2002; Kamboures  
177 et al., 2008; Meusel et al., 2018; Mushinski et al., 2019; Purchase et al., 2023; Song et al., 2023;  
178 Sörgel et al., 2015; Wang et al., 2021). The same is true for direct emissions of  $\text{NO}_2$  from soils,

179 where evidence remains limited and the uncertainty is high (Huber et al., 2024; Zörner et al., 2016).  
180 Only a handful of field studies have directly measured soil NO<sub>2</sub> measurements, making it  
181 challenging to infer how much NO<sub>2</sub> might be emitted from soils. For example, Gong et al. (2025)  
182 estimate that fertilizer-induced soil NO<sub>x</sub> emissions contribute 0.84–2.20 Tg N yr<sup>-1</sup> globally, with  
183 uncertainty ies partly due to the a lack of NO<sub>2</sub> measurements. Their modelling indicates that  
184 suggests this underrepresentation likely leads to underestimation of summertime ozone  
185 enhancements (by 0.3–3.3 ppbv) in agricultural hotspot regions, and has been which has been  
186 implemented further in other atmospheric models (Ha et al., 2023; Tian et al., 2024). This N<sub>r</sub>  
187 exchange is intricately linked to in agricultural practices, as regions from excessive nitrogen inputs  
188 (e.g., fertilizers) are prime targets for chamber methodologies, as they amplify emissions of all N<sub>r</sub>  
189 species, namely NO<sub>x</sub>, N<sub>2</sub>O, HONO, and NH<sub>3</sub> from the impacted soils (Degaspari et al., 2020;  
190 Huber et al., 2020; Manco et al., 2025). These issues highlight the need for more direct soil NO<sub>2</sub>  
191 and HONO measurements, as well as simultaneous constraints on the entire N<sub>r</sub> suite being  
192 exchanged from soils.

193 Automated dynamic chambers deployed in situ for field observations and performing controlled  
194 experiments would provide a platform for capturing the magnitude, direction, and temporal  
195 variability of N<sub>r</sub> species or physical variables via the headspace gas or soil amendments while  
196 retaining soils in an intact state (Aneja et al., 2006). Thus, establishing an accessible dynamic  
197 chamber method for N<sub>r</sub> flux measurements is desirable. However, such a platform needs to  
198 undergo extensive validation to reduce flux bias from challenging N<sub>r</sub> species such as NH<sub>3</sub>. This  
199 important and necessary first step will allow a wider global study of surface-atmosphere N<sub>r</sub>  
200 exchange processes. One of the best existing examples to date of automated dynamic chamber  
201 design for N<sub>r</sub> measurements, is the custom-built system from Pape et al. (2009) who measured NO,  
202 NO<sub>2</sub>, and O<sub>3</sub> to deploy an unattended array of six samplers with destructive gas analyzers. In their  
203 system, a reference chamber was used to characterize system surface effects, while using a large  
204 volume flow through the headspace during chamber closure periods to quantify fluxes on the  
205 assumption that ambient levels were not dramatically changing (e.g. due to nearby point sources).  
206 This work synthesized many advantages from similar designs to study soil- and plant-atmosphere  
207 interactions, but the technique remains accessible only to researchers with in-house engineering  
208 design and fabrication facilities. In the intervening years, dynamic chambers for GHG fluxes have  
209 become widely commercialized to improve measurement capacity compared to static chamber

determinations and to make flux observations more accessible compared to conducting eddy covariance measurements.

Here, we bridge several gaps to link the atmospheric GHG and  $N_r$  flux communities to yield a dynamic flux system for  $CO_2$ ,  $CH_4$ ,  $N_2O$ ,  $NO$ ,  $NO_2$ ,  $HONO$ , and  $NH_3$ . First, we modify commercial dynamic chambers with large volume (72 L) and footprint ( $0.21\text{ m}^2$ ) originally designed for trace GHG flux measurements to make them suitable for quantifying the most prevalent  $N_r$  gas exchange fluxes at surface-atmosphere interfaces, meaning the apparatus is more widely available to the atmospheric community. FirstNext, we implement surface and hardware modifications to adapt the commercial chambers to minimize gas adsorption and transformations, so that more reactive gases such as  $HONO$  and  $NH_3$  can be added to the  $N_r$  flux analyte suite. We systematically characterized the transfer of both GHGs and  $N_r$  species by calculating fill and empty rates, transformed to time constants, to identify any surface interactions and/or transformations on the chamber surfaces. We then applied our modified commercial dynamic chambers to make flux measurements by equipping them with destructive gas analyzers for  $HONO$  and  $NO_x$  and a cavity ring-down spectrometer (Picarro G2509) for  $NH_3$ ,  $N_2O$ ,  $CO_2$ , and  $CH_4$  in lab experiments, or with a fully automated dual-chamber approach under field conditions through a pilot study with 30 minute closures to obtain a sufficient number of measurements to detect relevant fluxes with standard gas analyzers. Fluxes during the pilot study were assessed by rate of change determinations during closure periods and bias minimization through a mass balance approach during the pilot study to in a real agricultural field is demonstrated system capabilities for several  $N_r$  gases in an agricultural field. This community-accessible approach addresses key needs in by allowing more researchers to measure  $N_r$  exchange at the surface-atmosphere interface, with the capacity-added benefit over past systems to monitor fluxes of all species simultaneously with at least hourly time resolution when using gas analyzers with one-minute measurement frequencies.

## 2. Materials and methods

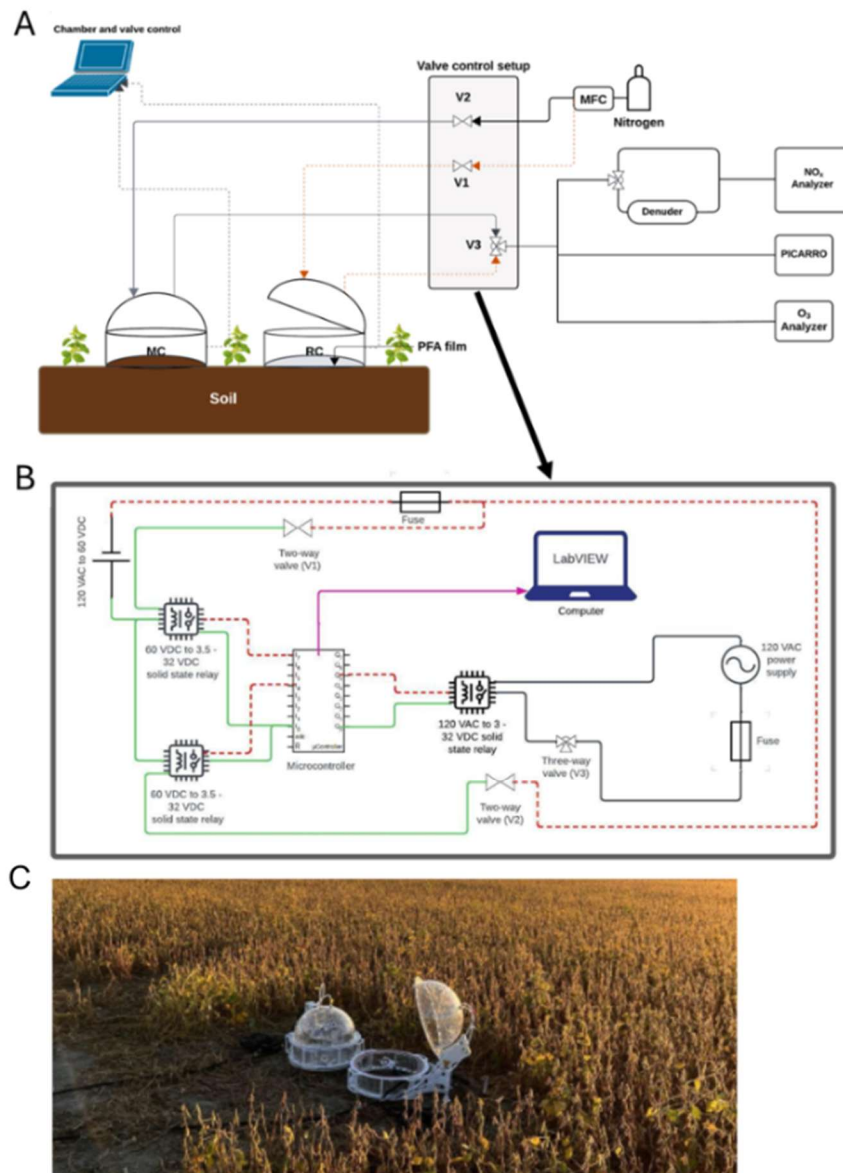
### 2.1 Dynamic Chambers for Field Fluxes

#### 2.1.1 Description of the custom-modified dual-dynamic chambers system for fluxes measurement

239 The dynamic chamber system ~~is composed of~~uses two identical commercially available ~~chambers~~  
240 ~~units~~ (eosAC-LT Eosense, Dartmouth, NS). These are modified and coupled to programmable  
241 valves that control ~~the delivery of~~ sample gases ~~delivered~~ to a suite of instrumentation (Figure 1)  
242 ~~to perform reactive gas flux measurements.~~

243 The dynamic chambers ~~are~~, constructed ~~out of with~~ transparent polyacrylate walls and lids, ~~have~~  
244 ~~with~~ an internal volume of 0.072 m<sup>3</sup> (72 L) and ~~can enclose a~~bottom surface area of 0.21 m<sup>2</sup>. When  
245 used on soils ~~surfaces~~, the chambers are secured with ~~embedded soil~~ collars ~~using and~~ custom-  
246 made polytetrafluorethylene (PTFE) rings (Figure S1) ~~and can be similarly affixed to impervious~~  
247 ~~surfaces~~. A built-in fan ensures uniform distribution of gases inside the chamber. ~~Ambient air~~  
248 temperature is measured inside ~~the chamber~~ from the fan arm, ~~whereas~~ pressure ~~sensors are is~~  
249 ~~embedded in from~~ the control box outside the chamber. ~~Each chamber has and~~ two auxiliary  
250 ports, one internally and ~~the other~~ externally, ~~capable of collecting measurements of can collect~~  
251 environmental properties such as relative humidity (RH), photosynthetically active radiation  
252 (PAR), soil temperature, and/or soil volumetric water content (VWC).

253 For a N<sub>r</sub> sampling approach, one ~~of the two~~ chambers is used as the measurement (MC) ~~to~~  
254 ~~enclose from~~ an experimental surface while the second ~~chamber acts as is~~ a reference (RC) ~~by~~  
255 ~~being~~ sealed at the bottom with a 51 μm (0.002") film of 0.002" perfluoroalkoxy alkane (PFA;  
256 McMasterr-Carr®, PN: 84955K24). The inert PFA film is held in place between the chamber  
257 collar and our custom-made PTFE rings (Figure S1). ~~This modification provides~~The RC acts as a  
258 negative control ~~observation~~ for physical interactions ~~on apparatus surfaces and/or~~ associated  
259 chemistry ~~when sampling of~~ reactive gases ~~on chamber and gas transfer line surfaces~~. The use of  
260 the RC, therefore, is to facilitate the correction of surface-mediated effects, reactions, and  
261 reduction of bias when determining quantitative flux values.



262

263 **Figure 1.** (A) Schematic of the dynamic chamber system to measure  $N_r$  and GHG fluxes. The  
 264 components of the system include: the chambers, a dilution gas source (nitrogen ( $N_2$ ) or zero air  
 265 (ZA)), solenoid valve control, gas transfer lines, and gas analyzers. Grey lines indicate dilution gas  
 266 flow from a source (e.g. cylinder) to the measurement chamber (MC) and air sampled from the  
 267 MC to the analyzers. The dashed orange line represents these same flows relative to the reference  
 268 chamber (RC). Communication lines from the chambers to a computer for automated control and  
 269 ancillary sensor data collection using the chamber software (eoslink-AC; blue dotted lines). (B)  
 270 The valve control setup for flow control in the complete dynamic chamber system, illustrating  
 271 electrical components and lines needed for full automation using LabVIEW. It includes two 24 V  
 272 DC two-way valves (V1 and V2; red dashed lines for negative electric potential, green for  
 273 positive), with power supplied by solid-state relays, and a three-way 120 V AC valve (V3; black  
 274 lines) with a power supply and another solid-state relay. The purple arrow represents the USB data  
 275 acquisition connection from the microcontroller to a computer running the LabVIEW VI for valve

276 control. (C) Field deployment of the dynamic chamber system, where the closed chamber is the  
277 reference and the open chamber functions as the measurement chamber.

278 ~~Reactive gases require continuous correction for surface interactions and/or transformations as~~  
279 ~~ambient levels are dynamic, so t~~The RC component of this system is designed to continuously  
280 ~~baseline~~ the physical interactions and chemistry happening on its surfaces both before and after  
281 quantifying reactive gas fluxes with the MC. Thus, the measurements are taken every 30 minutes,  
282 where one chamber is closed for gas analysis while the other is open to the ambient atmosphere.  
283 The sampling time interval was determined based on i) obtaining enough measurements at 1-  
284 minute time resolution to perform a reliable accumulation or loss linear regression, and ii) an  
285 ability to readily detect the lower limit of the range of field HONO flux values reported in the  
286 literature to ensure detection in its application during our pilot field study (see Section 2.76). For  
287 the first criterion, this also includes an exclusion of the first and last few measurements (3 to 5) to  
288 allow for complete gas replacement between the chamber lines and the analyzers, and disruption  
289 of the sealed environment as the chamber cycle alternates, respectively. The resulting accumulated  
290 mixing ratios of HONO in the chamber are well above the 1.4 parts per billion (ppbv) mixing ratio  
291 detection limit (LOD) of even a modified NO<sub>x</sub> analyzer, such that it can be used for measuring  
292 HONO (Crilley et al., 2023; Lao et al., 2020; Zhou et al., 2018; Nodeh-Farahani et al., 2021).

293 Headspace recirculation to facilitate analyte mixing ratio ~~enhancements~~ accumulation or depletion  
294 for non-destructive spectroscopic GHG analysis is a common measurement approach to decrease  
295 flux observation times. Reactive nitrogen measurements, in contrast, are typically destructive  
296 techniques that change the identity of the target analyte in the act of quantifying its abundance. To  
297 interface with such instruments, the sampled air needs to be replaced (Linde Canada Plc, PN: NI  
298 LC250-230) to balance the flow demand in a closed chamber. This balance is delicate even when  
299 using mass flow controllers (MFCs) on both incoming and outgoing flows, and the best solution  
300 we identified is to provide a slight overflow that takes advantage of the chamber design to vent  
301 excess pressure through a short length (~15 cm) of 1/8" ID (3.175 mm), 1/4" OD (6.35 mm) tubing  
302 that keeps the internal pressure equivalent to ambient. The flow differential between make-up gas  
303 and sampling is roughly 400 cm<sup>3</sup> min<sup>-1</sup>. Such a supply of make-up gas was explored across a range  
304 of potential flow rates when using destructive gas analyzers (e.g., three instruments each sampling  
305 at 1 - 4 standard litres per minute, L min<sup>-1</sup>) to find that 6 L min<sup>-1</sup> is the upper limit of flow-through

306 where the chamber pressure is not substantially perturbed from ambient and the chamber lid retains  
307 its seal.

308 In the field, a flux measurement cycle begins with closing the RC while the MC is open. At defined  
309 intervals, they alternate their open-closed states. Flows of make-up gas to each chamber are  
310 modulated with a pair of two-way solenoid valves. When sampling from the RC, one valve (V1;  
311 Figure 1) is open to permit make-up gas flow while the other valve (V2) is closed to prevent the  
312 flow from being directed to the MC. On the sampling lines, a three-way solenoid valve (V3)  
313 alternates to guide flow from whichever is closed to the suite of gas analyzers. ~~In this work, the  
314 gas instrumentation included a modified NO<sub>x</sub> chemiluminescent analyzer (NO, NO<sub>2</sub>, and HONO),  
315 a greenhouse gas and ammonia cavity ringdown spectrometer (H<sub>2</sub>O, CO<sub>2</sub>, CH<sub>4</sub>, N<sub>2</sub>O, NH<sub>3</sub>), and a  
316 UV-absorption ozone (O<sub>3</sub>) analyzer.~~ All instrument and operational details are provided in Section  
317 **2.1.3.**  
318

### 319 **2.1.2 Automated controls: system, data collection and processing**

320 The chamber eosLink-AC software (Eosense Inc., Dartmouth, NS) is used to define the duration  
321 of chamber opening and closing cycles, and logs chamber temperature, pressure, and auxiliary  
322 sensor data associated with a given eosAC-LT chamber. Each chamber requires a 12 V DC power  
323 supply connected by USB to a laptop through a weatherproof communication cable, controlling  
324 the chamber lid and data transfer.

325 When a chamber cycle begins, a text file is generated and includes measurement time elapsed,  
326 chamber lid status, chamber temperature and pressure, and auxiliary sensor data. This data file is  
327 updated at least once every 10 seconds, varying between 2-8 second intervals, which we average  
328 onto a 1-minute time base to match measurements from the slowest gas analyzers. The solenoid  
329 valves are modulated by the electrical circuit shown in Figure 1B. Automation is facilitated by a  
330 microcontroller (NI-6509i, National Instruments) programmed with a custom-scripted LabVIEW  
331 VI (LabVIEW version 2020). ~~The VI controls valve states (i.e., for V1, V2, and V3) and  
332 synchronization of the valve switching with respect to each chamber opening and closing cycle as  
333 described in the previous section. The VI also generates its own text file containing valve  
334 open/close state and a timestamp to be used for data processing. A custom R script (R-Studio  
335 v3.0.1) was developed to process the data file generated by eoslink-AC, data from all reactive gas~~

336 ~~analyzers, and the VI.~~ Further design information and full details of this sampling strategy and  
337 script can be found in Section S1 of the supporting information, and is available on the GitHub  
338 repository alongside our VI (<https://github.com/fjs-vdblelab/fluxchamber.git>).  
339

### 340 2.1.3 Reactive and greenhouse gas instrumentation for flux measurements

341 The mixing ratios of NO and NO<sub>2</sub> were measured using a commercially available  
342 chemiluminescent NO<sub>x</sub> analyzer (EC 9841, American Ecotech, Warren, RI). The calculated LOD  
343 determined from sampling dry zero air was 0.84 ppbv, 0.67 ppbv and 1.07 ppbv for NO, NO<sub>x</sub>, and  
344 NO<sub>2</sub> (or HONO when using the denuder as described below), respectively. The instrument has an  
345 operating range of 0 – 20 parts-per-million by volume (ppmv), a sample flow rate of 0.5 L min<sup>-1</sup>,  
346 and reports measurements at a time resolution of 1 minute. To quantify NO<sub>2</sub>, it is reduced to NO  
347 on a heated molybdenum catalyst (325 °C). To prevent interferences reported by others from  
348 atmospherically-relevant acidic species in this system (e.g. HONO, HNO<sub>3</sub>, and N<sub>2</sub>O<sub>5</sub>) the sampled  
349 air from the chambers during field experiments was passed through a sodium carbonate (Na<sub>2</sub>CO<sub>3</sub>)  
350 coated annular denuder to reduce bias in the NO<sub>2</sub> measurement, as these species and other  
351 components of NO<sub>y</sub> (e.g. peroxyacetyl nitrate; PAN) may also be reduced to NO (Villena et al.,  
352 2012). The Na<sub>2</sub>CO<sub>3</sub> denuder was prepared according to the EPA Compendium Method IO-4.2  
353 (Winberry Jr et al, EPA, 1999) to remove atmospheric acids by reactive uptake to the basic coating.  
354 As part of our controlled laboratory and pilot field study experiments, this denuder was also used  
355 to selectively measure HONO by scrubbing this target gas for a specified period, but would include  
356 the other known interferences. If this term is depositing, it could include the other NO<sub>y</sub> detected  
357 by the same conversion mechanism, but if emitting we expect it to be dominated by HONO.  
358 Ideally, a platform like time-of-flight chemical ionization mass spectrometry (ToF-CIMS) would  
359 be used for disambiguation but was not available at this time.

360  
361 A commercial O<sub>3</sub> analyzer (Serinus 10, American Ecotech, Warren, RI) was used to measure  
362 mixing ratios, quantify O<sub>3</sub> loss to surfaces, and constrain the reaction of O<sub>3</sub> with NO to form NO<sub>2</sub>  
363 in the pilot study sampling. This analyzer employs a non-dispersive UV absorption cell to quantify  
364 O<sub>3</sub> in the sampled air. The calculated LOD from sampling zero air is 0.95 ppbv at 1 minute time

365 resolution, with an operating range of 0 to 20 ppmv and a sampling flow rate of 0.5 L min<sup>-1</sup>. Quality  
366 control procedures for the NO<sub>x</sub> and O<sub>3</sub> instruments can be found in Section S2.

367  
368 The mixing ratios of ~~greenhouse gases (the GHGs) N<sub>2</sub>O, CH<sub>4</sub>, CO<sub>2</sub>, H<sub>2</sub>O, and ammonia (NH<sub>3</sub>)~~  
369 sampled from the automated chamber system were measured using a Picarro G2509 which uses  
370 cavity ring down spectroscopy (CRDS) ~~to simultaneously measure nitrous oxide (N<sub>2</sub>O), methane~~  
371 ~~(CH<sub>4</sub>), carbon dioxide (CO<sub>2</sub>), and NH<sub>3</sub> at ppbv levels, as well as water (H<sub>2</sub>O) vapor at ppmv~~. The  
372 analyzer has a time response of ~ 8 seconds for N<sub>2</sub>O, CH<sub>4</sub>, CO<sub>2</sub>, H<sub>2</sub>O and < 2 min for NH<sub>3</sub>. The  
373 Picarro G2509 ~~instrument~~ was used for ~~measurements during both~~ the lab experiments and the  
374 pilot field study ~~campaign~~. The customized version of the instrument ~~had a sampling flow rate of~~  
375 sampled at ~0.23 L min<sup>-1</sup>, and was equipped with an inlet filter. To minimize adsorption and  
376 chemical interactions of NH<sub>3</sub> on instrument surfaces, stainless steel gas handling components,  
377 including the inlet bulkhead, were replaced with PFA counterparts. The instrument cavity material  
378 was treated with a SilcoNert® coating by the manufacturer. We did not observe changes in its  
379 performance for the measured gases when operated according to the manufacturer guidelines. The  
380 Picarro G2509 analyzer utilizes cavity ringdown spectroscopy to quantify the mixing ratio of the  
381 target analytes, and determination means a full span calibration is not a regular necessity to be  
382 performed regularly. Despite this, we validated its calibration and performed quality control  
383 checks in ~~our the~~ laboratory to ensure the accuracy and stability of the analyzer for all aspects of  
384 this work presented below (Section S2).

385

## 386 2.2 Chamber modifications to minimize NO<sub>2</sub> reactions on chamber surfaces

387 To transfer reactive gases through these chambers, interactions with surfaces need to be limited at  
388 all points of potential adsorptive or reactive losses. The custom-made base plate, ~~consisting of two~~  
389 ~~rings made from PTFE sheets with a PFA film pinned between them~~ (Figure S1) was used to assess  
390 gas interactions on the commercial chamber surfaces ~~and acts as the RC for field measurements~~.  
391 ~~From this starting point, commercial components were and identification ed for of parts for~~  
392 replacement. First, the gas inlet and outlet push-to-connect fittings in the original ~~chamber~~  
393 configuration have plastic grips, with an internal component made of brass, which is informally  
394 known in the atmospheric chemistry community to have strong interactions with nitrogen oxides.

395 These ~~brass fittings~~ were replaced with PTFE Swagelok<sup>®</sup> bulkhead fittings (PN: T-400-1-4) ~~as~~  
396 ~~shown in;~~ Figure S2). ~~Similarly~~ ~~Second~~, the polyacrylate wall and lid surfaces of the chambers  
397 ~~were considered, which are, by design, optically transparent to ensure PAR is transferred to any~~  
398 ~~contained plants or surfaces. To retain this visible radiation transparency, while also making the~~  
399 ~~surface more inert, we applied~~ had the 51  $\mu\text{m}$  (0.002") thin-PFA film ~~applied~~ to the inner surfaces  
400 using double-sided tape ~~to retain actinic transparency and PAR transfer to contained plants and~~  
401 ~~surfaces~~ (Figure S2).

### 403 2.3 Chamber modification validation using greenhouse and reactive gases

404 ~~The design of the chambers by the manufacturer is to transfer GHGs to non-destructive gas~~  
405 ~~analyzers by recirculating the headspace.~~ Before and after our modifications, we had to ensure that  
406 the non-reactive GHG ~~transfers were still transferred through the system was unchanged.~~ In  
407 addition, ~~we to~~ challenged ~~tests for~~ the transfer of  $\text{N}_r$  gases ~~by controlled gas delivery.~~ It was  
408 expected that  $\text{N}_r$  gases ~~could undergo would~~ interactions and/or ~~chemical transformations~~ ~~react on~~  
409 ~~the chamber surfaces and 15 m of standard sampling tubing that would differ between the~~  
410 ~~unmodified and modified variants.~~ Determining the time constants of fill (E1, E2) and decay (E3,  
411 E4) of these ~~interactive and/or~~ reactive gases in the chamber ~~system~~ allowed us to contrast their  
412 behaviour against that expected from a modelled theoretical inert trace gas in our system (Figure  
413 2). Equations used to model mass transfer in our chambers were derived from Pape et al. (2009).  
414 The resulting accumulation curve was modelled by the theoretical function:

$$415 \quad \mu_{\text{fill}}(t) = 1 - e^{-\left(\frac{t}{\tau_{\text{fill}}}\right)} \quad (\text{E1})$$

$$416 \quad \tau_{\text{fill}} = V/Q_{\text{fill}} \quad (\text{E2})$$

417 Where  $\mu_{\text{fill}}$  represents the normalized mixing ratio of the gas in the chamber at the time  $t$  (min)  
418 after closing of the chamber compared to the maximum mixing ratio within the measurement cycle,  
419  $\tau_{\text{fill}}$  is theoretical accumulation timescale for transfer of an ideal inert gas (min),  $V$  is the volume  
420 of the chamber ( $0.072 \text{ m}^3$ ), and  $Q_{\text{fill}}$  represents the total experimental flow rate ( $2 \text{ L min}^{-1}$ ).  
421 Similarly, the theoretical decay curve when emptying the chamber can be obtained, where  $\tau_{\text{emp}}$  is  
422 the theoretical decay timescale for gas transfer (min) and  $Q_{\text{emp}}$  is again the total experimental flow  
423 rate ( $2 \text{ L min}^{-1}$ ).

424 
$$\mu_{emp}(t) = e^{\left(-\frac{t}{\tau_{emp}}\right)} \quad (\text{E3})$$

425 
$$\tau_{emp} = \frac{V}{Q_{emp}} \quad (\text{E4})$$

### 426 2.3.1 Instrumentation and materials for control experiments

427 Control experiments for the transmission of ~~inert-GHG~~ and ~~reactive-N<sub>r</sub>~~ gases were conducted by  
428 filling and emptying the chambers with known quantities ~~of the target gases~~ at mixing ratios  
429 relevant to the atmosphere, as well as ~~in~~ quantities ~~that would expected to~~ accumulate during real  
430 observations of modest ~~emission~~ fluxes (e.g. from a fertilized farm field). ~~Measurement of NO,~~  
431 ~~NO<sub>2</sub> and HONO was performed using the modified NO<sub>x</sub> analyzer, the GHGs (CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O)~~  
432 ~~and NH<sub>3</sub> with the Picarro G2509, and O<sub>3</sub> with the UV absorption instrument. All assessments~~  
433 ~~herein matched: the standard configuration of the chambers with all fittings, 15 m of ¼" O.D. PFA~~  
434 ~~gas transfer tubing, flow rates, valves and gas transfer lines to instrumentation, with~~  
435 ~~line/fitting/valve/instrumentation surfaces included.~~ Details of the generation of gas concentrations  
436 can be found in Section S3 of the SI.

### 437 2.3.2 Filling and emptying experiments with N<sub>r</sub>, O<sub>3</sub>, and GHGs

438 The positive control experiments filled the chambers in both modified and unmodified  
439 configurations, ~~where the and~~ time constants were calculated from the measurements ~~to compare~~  
440 ~~against a theoretical inert and perfectly transferred gas (i.e. lost by dilution and removal to~~  
441 ~~analyzers only).~~ In each filling experiment, the chamber was flushed with pure N<sub>2</sub> from a liquid  
442 N<sub>2</sub> dewar (Linde Canada, PN: NI LC250-20) until a stable baseline level of each gas mixing ratio  
443 was reached; typically, these were values at the analyzer detection limits. Then, a blend of GHGs  
444 or one of the N<sub>r</sub> analytes was delivered into the chamber ~~along~~ with N<sub>2</sub> ~~as the~~ dilution flow at a  
445 total flow rate of 2 L min<sup>-1</sup>, ~~and this which~~ was then sampled at ~~a total of~~ 1.8 L min<sup>-1</sup> by the  
446 analyzers (Figure S4). The gases were added to the chamber from their respective calibration  
447 sources until the observed concentration (C) reached the known value being delivered (C<sub>0</sub>), within  
448 error. Since different mixing ratios of the gases were added for these control experiments, the use  
449 of normalized concentrations was necessary to facilitate data analysis and visualization. Where  
450 surface interactions could be identified (e.g. for NH<sub>3</sub>), the role of surfaces versus air exchange  
451 ~~whas been~~ explored using double exponential fits (see ES-1 and ES-2 in Section S3) (Crilley et  
452 al., 2023; Ellis et al., 2010; Moravek et al., 2019). The gases were emptied back to the initial

453 baseline level ~~observed with clean dilution gas~~ before starting the next replicate or a new ~~filling~~  
454 experiment with a different target gas. Time constants for filling and emptying were ~~then~~  
455 determined by fitting the ~~experimental~~ observations in Igor Pro8 (Wavemetrics, Portland, OR,  
456 US).

457 Similarly, the Eosense eosMX multiplexer is designed to coordinate chamber flux measurements  
458 using eosAC chambers with non-destructive analyzers, such that headspace can be recirculated  
459 while the chambers are closed. One of these devices was ~~also~~ characterized ~~for the transmission of~~  
460 ~~GHGs and N<sub>r</sub> gases, alongside following~~ similar modifications. This system is appealing as it has  
461 the eosLink-MX software (Eosense, V1.9.07), which is used for communication, scheduling  
462 actions, and logging peripheral data from all connected eosAC chambers. It features dedicated  
463 chamber tubing inlets and outlets, along with a COMM port supporting up to 12 eosAC chambers.  
464 Each chamber channel includes two Swagelok gas fittings for transporting gases to analyzers and  
465 for either recirculating, or in the case of N<sub>r</sub> measurements, supplying a ~~purging clean dilution~~ gas  
466 to the chamber headspace.

467 To optimize the performance for N<sub>r</sub> species, the original stainless-steel (SS) Swagelok fittings and  
468 solenoid valves were compared against replacement PFA tube fittings, bulkhead unions  
469 (Swagelok, PFA-420-61), and a PTFE 3-way valve (Clippard, NR1-2-12-G2). A 2 L min<sup>-1</sup> flow  
470 of dry zero air containing the target compounds was passed through either the SS valve with SS  
471 fittings or the PTFE valve with PFA fittings for 30 minutes each. The flow was measured before  
472 and after the valves to ensure the setup was free of leaks. **The ratio of the transferred gas amount**  
473 **to the nominal was used to identify impacts of surface interactions on quantitative transmission to**  
474 **downstream gas analyzers. Further details are presented in Sections 2.7 and S4.**

#### 475 **2.4. Characterization of NO<sub>2</sub> and O<sub>3</sub> loss on chamber surfaces**

476 All losses of NO<sub>2</sub> in the chamber were characterized by the addition of known NO<sub>2</sub> mixing ratios  
477 (5-10 ppbv) to the chamber under ~~a variety of~~ relevant relative humidity (RH) conditions (45-85%  
478 RH; Figure S5; Section S5). The mixing ratios of NO<sub>2</sub> and RH values selected for these  
479 experiments are representative of the ambient atmosphere in urban areas (Toronto North Station,  
480 ECCC), where the chamber system was envisioned to be deployed. –The experiments were  
481 performed progressively and in triplicate with 5 ppbv of NO<sub>2</sub> and 85% RH, starting from the  
482 unmodified configuration of the chamber, replacement of fittings, and covering the inner surface

483 with PFA film to quantify their efficacy in minimizing NO<sub>2</sub> loss and/or transformation on chamber  
484 surfaces. These were followed by varying NO<sub>2</sub> mixing ratios and RH to characterize the modified  
485 system.

~~486 Different RH values were produced inside the flux chamber by combining flows of dry zero air  
487 (ZA) and one saturated with water vapour by transiting a 500 mL Pyrex impinger filled with  
488 deionized water (e.g. equal 1 L min<sup>-1</sup> flows combine to an RH of 50%). A high-precision, research-  
489 grade humidity probe (HMP60, Vaisala Oyj, Finland; ± 3% at 0–90% RH, ± 5% at 90–100%  
490 RH) was connected to the chamber to confirm the set point by measurement in the chamber. The  
491 flow from an NO<sub>2</sub> calibration cylinder (Linde Canada; PN: NI NX5MC-AQ; 5.9 (±5%) ppm) was  
492 adjusted using an MFC and combined with a dilution flow of ZA to achieve mixing ratios of 5, 7  
493 and 10 ppbv. The total flow rate for these experiments was 2.0 L min<sup>-1</sup> and a vacuum pump (62.3  
494 L min<sup>-1</sup>, PN: UZ-07061-22, Gast Manufacturing Inc., Benton Harbor, MI, USA) with an MFC (10  
495 L min<sup>-1</sup>, PN: 1179C01314CR1BV, MKS instruments Inc, Andover, MA, US) was used to make  
496 up the sampling flow beyond the 0.5 L min<sup>-1</sup> of the NO<sub>x</sub> analyzer. The same flow difference stated  
497 above was maintained in the chamber, using the built-in vent.~~

498 Quantification of NO<sub>2</sub> and HONO in the air sampled from the chamber can be achieved using the  
499 alternating solenoid setup depicted in Figure S5. The sampled air is switched between two channels  
500 – one directly to the NO<sub>x</sub> analyzer and the other through a Na<sub>2</sub>CO<sub>3</sub>-coated denuder – modulated  
501 every 5 minutes by a three-way solenoid valve (Fluoroware Galtex 1/4" F-NPT 3-way solenoid  
502 valve, 115V, PN: 203-3414-415. Entegris Inc., MN, US). When the sampled air flows directly to  
503 the analyzer, the total mixing ratio of NO<sub>2</sub> and acidic NO<sub>y</sub> species, like HONO and HNO<sub>3</sub>, is  
504 measured and has been termed NO<sub>2</sub><sup>\*</sup> (Crilley et al., 2023; Lao et al., 2020; Zhou et al., 2019).  
505 When the flow is directed through a Na<sub>2</sub>CO<sub>3</sub> denuder, it selectively scrubs HONO and HNO<sub>3</sub>,  
506 leaving behind NO<sub>2</sub> (Possanzini et al., 1983). Under the controlled NO<sub>2</sub> composition used in our  
507 experiments, it is expected that HONO will be the only acid present in the sampled air, as such  
508 experimental systems have been thoroughly characterized by Finlayson-Pitts et al. (2003),<sup>5</sup> and  
509 HNO<sub>3</sub> is retained on the surfaces (R3) (Barney & Finlayson-Pitts, 2000; Huang et al., 2002;  
510 Kamboures et al., 2008). As a result, by the differential measurement of mixing ratios recorded in  
511 the two channels (E5) every 5 minutes, HONO can be quantified.

$$512 \text{ HONO} = \text{NO}_2^* - \text{NO}_2 \quad (\text{E5})$$

513 To quantify the amount of NO<sub>2</sub> lost to the chamber surface relative to the mixing ratio of NO<sub>2</sub>  
514 added to the chamber during an experiment, the NO<sub>2</sub> loss fraction ( $f_{\text{NO}_2}$ ) ~~is~~ was found to be  
515 informative for mass balance between the two processes. ~~Comparison of each chamber~~  
516 ~~modification on  $f_{\text{NO}_2}$  helps identify which changes are essential in minimizing losses. Experiments~~  
517 ~~using 5 ppbv of NO<sub>2</sub> at 85% RH added to the chamber were used to quantify  $f_{\text{NO}_2}$  in the unmodified~~  
518 ~~configuration (n=3), to derive an upper limit of irreversible loss to surfaces or chemical loss via~~  
519 ~~surface hydrolysis (R3). Following the materials characterization experiments, additions of~~  
520 ~~atmospherically relevant 5, 7 and 10 ppbv mixing ratios of NO<sub>2</sub> at three different RH values (45%,~~  
521 ~~65% and 85%) to the fully modified chamber were used to characterize and develop a correction~~  
522 ~~methodology for the surface loss. The mixing ratios of NO<sub>2</sub> and RH values selected for these~~  
523 ~~experiments are representative of the ambient atmosphere in urban areas (Toronto North Station,~~  
524 ~~ECCC), where the chamber system was envisioned to be deployed.~~

## 525 **2.5 Drivers of O<sub>3</sub> loss to chamber surfaces**

526 ~~Similarly, We~~ tested ~~how the~~ modifications and aging of the ~~interior-installed PFA chamber film~~  
527 ~~surfaces affected on~~ O<sub>3</sub> transfer. It is among the most sensitive/reactive species to transfer through  
528 this new system and is expected to drive reactive loss of NO when measuring N<sub>r</sub> fluxes. The  
529 fraction of O<sub>3</sub> lost to chamber surfaces was ~~first~~ quantified ~~from 150-250 ppbv~~ using a clean and  
530 unmodified chamber with PTFE bulkhead fittings replacing the push-to-connect brass fittings  
531 ~~(Section S5). Prior to O<sub>3</sub> addition, the chamber was flushed with ZA until the background level of~~  
532 ~~O<sub>3</sub> was at the instrument detection limits. Three mixing ratios of O<sub>3</sub> of 150, 200 and 250 ppb~~  
533 ~~generated by a GasCal 1100 dilution calibrator with integrated O<sub>3</sub> generator and photometer~~  
534 ~~(American Ecotech, Warren, RI) and added to the chamber for 60 mins or until a constant~~  
535 ~~concentration was reached. Two replicate runs of each addition level were performed.~~

536 The fraction of O<sub>3</sub> lost to chamber surfaces was ~~then~~ quantified in duplicate on a modified chamber  
537 with the interior chamber surfaces covered by brand new 0.002" thick PFA film ~~in a separate set~~  
538 ~~of experiments. The same mixing ratios of O<sub>3</sub> were delivered in duplicate to the modified chamber~~  
539 ~~operated for lab and field experiments, so that the PFA film had been attached to the inner surface~~  
540 ~~of the chamber and or one~~ exposed to ambient air for more than two years, with 15 days of  
541 continuous use in an agricultural field during our pilot study. ~~In all three experiments, the lost~~  
542 ~~fraction of O<sub>3</sub> was determined for each mixing ratio delivered.~~

543 **2.56 Proof of concept N<sub>r</sub> fluxes from agricultural soils**

544 **2.56.1 Soil sample N<sub>r</sub> emissions for lab experiments**

545 Randomized soil samples weighing 4-5 kg were collected into Ziploc<sup>®</sup> bags from an eight-plot  
546 grid established ~~for at~~ an agricultural field site in Lambton County, ON, Canada (43°09'36.0" N  
547 81°55'48.0" W). The samples were used to investigate emissions in the lab with and without the  
548 addition of fertilizers. Individual and pooled samples from the field plots were used. Bulk soil  
549 samples were prepared for lab-based chamber measurements by removing debris, roots and seeds,  
550 followed by oven drying at 35 °C for 24-72 hours on a stainless-steel mesh tray covered with  
551 aluminum foil, to prevent alteration of the microbial community from exposure to unrealistic  
552 temperature regimes. After drying, samples were stored in Ziploc<sup>®</sup> bags at room temperature until  
553 use.

554 Ultrapure water (18 MΩ·cm; Milli-Q<sup>®</sup>, Sigma-Aldrich, St, Louis, US) was added to  
555 approximately 350 g of a dry soil sample to achieve ~28% volumetric water content (VWC). The  
556 soil sample was loaded into the chamber on a foil-lined tray, and the water content was measured  
557 using a soil moisture probe inserted fully into the sample (TEROS 11, VWC range for mineral  
558 soils: 0.00 – 0.70 m<sup>3</sup>/m<sup>3</sup>; accuracy: ± 0.03 m<sup>3</sup>/m<sup>3</sup>; resolution: 0.001 m<sup>3</sup>/m<sup>3</sup>, METER Group Inc.,  
559 WA, USA). Zero air modified to 65% RH was delivered at 3.6 L min<sup>-1</sup> to the chamber headspace  
560 where the soil was contained. The soil started the drying process from a VWC of approximately  
561 25% with the flows held constant for around 4 days or until the VWC reached 15%. The chamber  
562 was ~~then sealed against intrusion from room air, and dry ZA was added to the chamber during the~~  
563 ~~to conduct the~~ drying ~~period~~ cycle while our ~~ZA-modified NO<sub>x</sub> analyzer measured NO, NO<sub>2</sub> and~~  
564 ~~HONO following our usual configuration, while and~~ the Picarro G2509 measured N<sub>2</sub>O and NH<sub>3</sub>  
565 fluxes. ~~Both instruments monitored u~~ Unamended soil samples and others fertilized with urea  
566 (CO(NH<sub>2</sub>)<sub>2</sub>), ammonium carbonate (AC, (NH<sub>4</sub>)<sub>2</sub>CO<sub>3</sub>), ammonium bicarbonate (ABC, NH<sub>4</sub>HCO<sub>3</sub>)  
567 at rates of 100 kg N ha<sup>-1</sup> were assessed. A similar experiment was conducted with ammonium  
568 nitrate (NH<sub>4</sub>NO<sub>3</sub>) at the same fertilizer rate using only the modified NO<sub>x</sub> analyzer. Soil VWC and  
569 headspace RH were recorded using auxiliary sensors within the chamber.

570 **2.56.2 Field deployment of automated dynamic N<sub>r</sub> chambers**

571 The RC and MC setup was deployed to make automated N<sub>r</sub> flux measurements from the same  
572 agricultural field as ~~the lab based soil samples described~~ in the prior section. The observations

573 took place in early September 2022 at the end of a soybean cropping season. A detailed description  
 574 of the campaign and its results is the subject of a separate work, so we provide a brief overview  
 575 here ~~relevant to demonstrating the utility and procedure of an RC/MC chamber pair to determine~~  
 576 ~~fluxes from field observations~~. The total measurement period was approximately two weeks in  
 577 duration to test ~~the system~~ performance ~~of the designed system~~. Generally, conditions were hot  
 578 and dry, without precipitation, and the soybeans surrounding the observed soils were undergoing  
 579 senescence during the measurement period. The chambers were deployed only on the soil, between  
 580 crop rows, and operated to quantify fluxes as outlined in Section S7. After an initial 7-day period  
 581 of observing baseline fluxes from the field, an experimental perturbation was conducted to  
 582 stimulate  $N_r$  emissions through the addition of an aqueous urea solution equivalent to 22 kg N ha<sup>-1</sup>  
 583 of fertilizer added by broadcast application, followed by washing into the soil by an equivalent ~~cy~~  
 584 of 2.5 cm (1”) of rain ~~depth~~. The modified NO<sub>x</sub> analyzer and the Picarro G2509 were used to  
 585 measure the  $N_r$  and GHG fluxes continuously across both periods.

## 586 2.76 Soil flux determination

587 The flux of a gas is the rate at which it is ~~released and/or~~ transferred across ~~a surface or~~ an interface  
 588 (e.g., soil to atmosphere) per unit area per unit time. Gas fluxes are of high interest in agriculture  
 589 as they give insight into the uptake or emission of N-bearing gases that may ~~lead to alter~~ fertilizing,  
 590 ~~or loss of fertilizing~~ effects, ~~respectively~~. ~~Fluxes~~ ~~They~~ are also important for assessing the state of  
 591 plants or soils at interfaces through metrics like primary productivity, in which case measurement  
 592 of a GHG like CO<sub>2</sub> provides the insight (Anthony & Silver, 2024; Li et al., 2016; Okiti et al.,  
 593 2025). **The RC captures environmental fluctuations such as temperature or pressure change and  
 594 directly observes the interactions of ambient gases with surfaces within the sampling setup (i.e.  
 595 chambers, gas transfer lines, valves, and analyzers), as well as tracking reactions, allowing for  
 596 corrections to every net flux ( $F_{net}$ ) measurement cycle (E6 for reactive gases and E7 for non-  
 597 reactive gases, as derived in Section S7).**

$$598 F_{net} = (\lambda) \cdot \left( \frac{V}{A} \left( \frac{\Delta C_m}{\Delta t_m} - \frac{\Delta C_r}{\Delta t_r} \right) + \frac{Q_{out}}{A} \left( \frac{\int_{t_{1m}}^{t_{2m}} c_m(t) dt}{\Delta t_m} - \frac{\int_{t_{1r}}^{t_{2r}} c_r(t) dt}{\Delta t_r} \right) - \frac{V}{A} \left( \frac{\int_{t_{1m}}^{t_{2m}} R_m(t) dt}{\Delta t_m} - \frac{\int_{t_{1r}}^{t_{2r}} R_r(t) dt}{\Delta t_r} \right) \right)$$

599 E6

600 Where  $V$  is the volume of the chamber ( $\text{m}^3$ ),  $A$  is the surface area ( $\text{m}^2$ ) enclosed by the chamber  
601 and governing the gas flux;  $Q_{out}$  is the volumetric flow rate of air exiting the chamber ( $\text{m}^3 \text{ s}^{-1}$ ;  
602  $c_m(t)$  and  $c_r(t)$  are target gas concentrations within the [measurement-MC](#) and [reference-chamber](#)  
603 [RC](#) ( $\text{mol m}^{-3}$ ), respectively;  $\frac{\Delta C_m}{\Delta t_m}$  and  $\frac{\Delta C_r}{\Delta t_r}$  represent their corresponding rates of change ( $\text{mol m}^{-3} \text{ s}^{-1}$ ); and  $F_{net}$  is the resulting net gas flux per unit area ( $\text{mol m}^{-2} \text{ s}^{-1}$ ). The terms  $R_m$  and  $R_r$  denote the  
604 instantaneous chemical production or loss rate expressed in units of  $\text{mol m}^{-3} \text{ s}^{-1}$  for consistency.  
605 **The dimensionless attenuation factor  $\lambda$  is required to correct for interactions of reactive gases with  
606 surfaces. Such surface interactions, which are particularly strong for gases like  $\text{NH}_3$ , significantly  
607 reduce the measured rate of concentration change within the closed chamber (Figure 4). Thus,  $\lambda$  is  
608 derived as the ratio between a theoretical unattenuated gas (i.e. an inert GHG like  $\text{N}_2\text{O}$ ) and the  
609 target gas concentration from controlled deliveries integrated over the chamber closure interval as  
610 derived in Section S7. [This term has the surface effects from chambers, gas transfer lines, and  
611 analyzers embedded by definition and must be determined empirically for any altered  
612 configuration.](#) The attenuation correction removes bias from [surface-induced artifacts in flux  
613 estimates, so that more accurate soil-atmosphere exchange is reported.](#)**

$$615 \quad F_{net} = \lambda \cdot \frac{P_{air}}{R \cdot T} \cdot \left( \frac{V}{A} \left( \frac{\Delta X_m}{\Delta t_m} - \frac{\Delta X_r}{\Delta t_r} \right) + \frac{Q_{out}}{A} \left( \frac{\int_{t_{1m}}^{t_{2m}} X_m(t) dt}{\Delta t_m} - \frac{\int_{t_{1r}}^{t_{2r}} X_r(t) dt}{\Delta t_r} \right) \right)$$

616 E7

617 For inert gases [described in by E7](#), the fluxes can be based on mixing ratios, where  $X_m$  and  $X_r$  are  
618 the gas volumetric mixing ratios (mol X per mol air),  $P_{air}$  is the air pressure (Pa),  $T$  is the absolute  
619 temperature (K), and  $R$  is the universal gas constant ( $\text{J mol}^{-1} \text{ K}^{-1}$ ). By comparing the RC and MC  
620 observations, we isolate the effects of specific environmental conditions on  $N_r$  (E6) or GHG (E7)  
621 exchange fluxes, while accounting for surface effects and chemical transformations in the former.

## 622 3 Results

### 623 3.1 Determining time constants of reactive nitrogen and GHGs

624

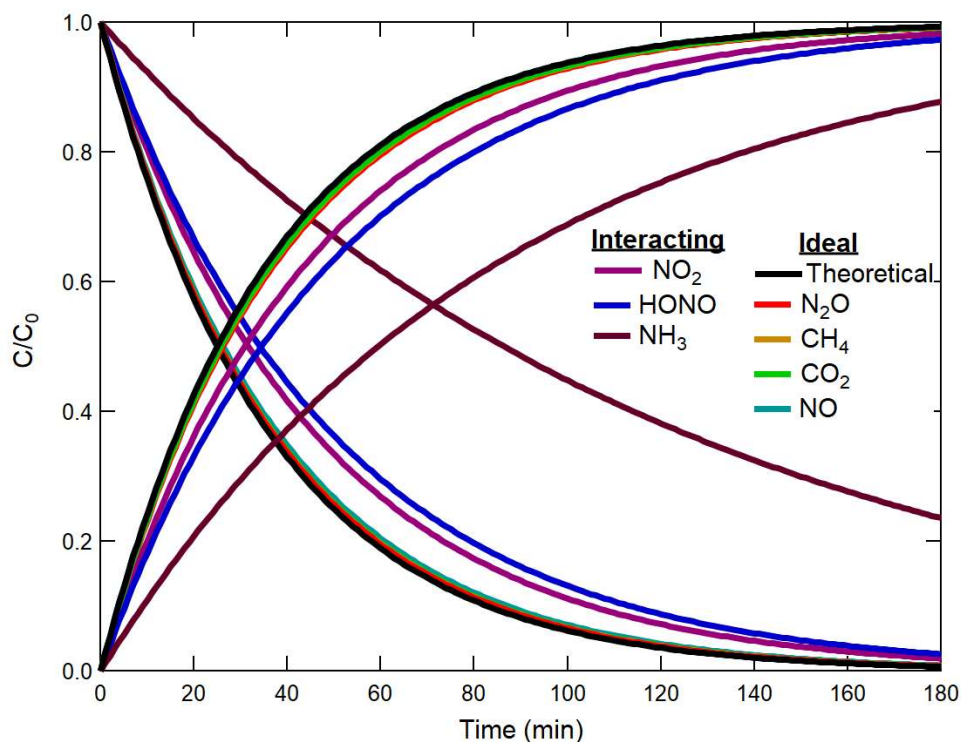
625 The time constants of both filling and emptying the chambers were calculated using concentrations  
626 normalized to their initial values, for each  $N_r$  gas and GHG. These were used to quantify gas  
627 transfer times through the chambers and to confirm performance relative to theory. Where

628 departures were identified, we quantified the extent of surface interactions for the various target  
629 analytes so corrections for determining in situ fluxes could be implemented (Table 1, Figure 2).

### 630 3.1.1 Time constants of greenhouse gases (GHGs)

631 Determination of the GHG time constants allows benchmarking of the chamber performance prior  
632 to and after modifications, ~~which then allows us to contrast the behaviour of transferred reactive~~  
633 ~~gases through the chambers.~~ In both ~~the modified and unmodified chambers, configurations~~ the  
634 greenhouse gases CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O were anticipated to behave as non-reactive trace gases with  
635 little to no physical interactions on chamber surfaces, nor to undergo chemical transformations.

636 The theoretical fill and empty rates for the chambers with a flow rate of 2 L min<sup>-1</sup> are 36 min. The  
637 average measured time constants of filling ~~with the greenhouse gases for~~ CH<sub>4</sub>, CO<sub>2</sub>, and N<sub>2</sub>O in  
638 the unmodified chamber were ~~found to be~~ 37 ± 1 min, 37 ± 2 min, and 37 ± 1 min, respectively  
639 (Table 1). During ~~the emptying phase,~~ ~~the~~ ~~average time constants for CH<sub>4</sub>, CO<sub>2</sub>, and N<sub>2</sub>O~~ were  
640 37 ± 1 min, 37 ± 1 min, and 38 ± 2 min, respectively. These measurements are not different from  
641 theory within the limits of experimental accuracy, ~~demonstrating that GHGs do not deviate from~~  
642 ~~the values modelled for an ideal trace gas~~ (Figure 2). ~~In either experiment, the time constants did~~  
643 ~~not change in the presence of the chamber modifications.~~ Since the GHGs are effectively  
644 transferred through both the modified and unmodified configurations of the chamber, the baseline  
645 performance of the chambers was not affected by the hardware modifications. Therefore,  
646 comparison of these results with the time constants of N<sub>r</sub> gases provides a description of their  
647 interaction or transformation processes on the modified chamber surfaces.



648

649 **Figure 2.** Addition of reactive nitrogen gases and greenhouse gases to a modified dynamic  
 650 chamber. For clarity, the coloured traces show the fitting curves corresponding to the response  
 651 time in concentration normalized to the delivered value of each gas while filling and emptying the  
 652 chamber. The black trace corresponds to a perfect non-reactive transfer of an ideal trace gas based  
 653 on volume transfer in the chamber only. Note that the N<sub>2</sub>O, CH<sub>4</sub>, CO<sub>2</sub> and NO fill and empty traces  
 654 all overlap with the theoretical fill and empty curves.

655

### 656 3.1.2 Time constants of reactive nitrogen gases

657 In the modified chamber, the time constant of filling or emptying with NO was  $38 \pm 1$  min. The  
 658 obtained NO time constants are similar to GHGs, as NO is not expected to have strong surface  
 659 interactions. The slower time response of the NO<sub>2</sub>, HONO, and NH<sub>3</sub> measurements is determined  
 660 by two processes: (1) the exchange of the sample air volume in the inlet-gas transfer lines and the  
 661 chamber, and (2) the adsorption and desorption of the gas onto and from the chamber-surfaces  
 662 (Whitehead et al., 2008). Increases in the extent of surface interactions were observed for followed  
 663 the increasingly polarity, reactivity, and/or ionizability of gases in the order of NO<sub>2</sub>, then HONO,  
 664 and most for NH<sub>3</sub> (Figure 2). The highest change in the time constants between the modified and  
 665 unmodified configurations was  $2.6 \pm 2.5$  min, observed for HONO, with smaller improvements

666 observed for NO<sub>2</sub> and NH<sub>3</sub> during the filling process. Improved time constants when emptying  
667 reactive nitrogen from the chambers had similar trends (Table 1).

668 Overall, the deviations for these gases to longer time constants than those expected from an inert  
669 molecule can be attributed to their polar, ionizable, and/or reactive nature. For example, NO<sub>2</sub> is  
670 lost more readily than NO, possibly through its known reaction on surfaces to make HONO and  
671 HNO<sub>3</sub>, being lost itself in the process (Finlayson-Pitts et al., 2003). It also has higher water  
672 solubility than NO, but lower than for HONO or NH<sub>3</sub>. Similarly, a decrease in transmission  
673 efficiency for HONO could be explained by its weakly acidic nature (pKa = 3.16) (Silva et al.,  
674 2006) and solubility in water (Henry's law constant = 0.48 mol/m<sup>3</sup> Pa; (Schwartz & White, 1981)  
675 that facilitate partitioning and dissociation in surface water films, which could generate non-  
676 volatile nitrite (NO<sub>2</sub><sup>-</sup>) on chamber surfaces. This chemistry will slow the transfer of gas-phase  
677 HONO through the chambers, as the NO<sub>2</sub><sup>-</sup> would need to protonate before being lost as neutral  
678 HONO when repartitioning to the gas phase (R4, R5). Finally, the most delayed transmission rate  
679 is for NH<sub>3</sub>, likely because it undergoes strong inter-molecule interactions and ionization on the  
680 chamber surfaces and/or any interfacial water (Henry's law constant = 5.9x10<sup>-1</sup> mol/m<sup>3</sup> Pa (Orkin  
681 et al., 2011); pKa = 9.25 (Lide, 2009), as well as on tubing surfaces and potentially partitioning  
682 into the tubing before reaching the analyzer (Pagonis et al., 2017). The best improvement in the  
683 time constants between the modified and unmodified configurations was 2.6 ± 2.5 min for HONO,  
684 with smaller improvements observed for NO<sub>2</sub> and NH<sub>3</sub> during the filling process. Improved time  
685 constants when emptying reactive nitrogen from the chambers had similar trends (Table 1). As a  
686 result, increasing delays from NO through NH<sub>3</sub> exist in our N<sub>r</sub> gas suite due to increasingly stronger  
687 interactions with chamber surfaces and gas handling lines.

688 The determined surface impact-interaction values (D; Table 1; ES-2, Section S3) demonstrate an  
689 expected greater impact of surfaces when no reactive gas is present in the headspace prior to filling,  
690 and a lesser effect during emptying when as the exposed surface has been exposed and equilibrated  
691 with the analyte, which is commonly referred to as passivation. As a result, increasing delays from  
692 NO through NH<sub>3</sub> exist in our N<sub>r</sub> gas suite due to increasingly stronger interactions with chamber  
693 surfaces and gas handling lines. For NH<sub>3</sub> specifically, the fill has a D value of 89%, while during  
694 emptying it is only 23%. This is due to all sample handling surfaces not being passivated prior to  
695 filling, similar to our findings with NH<sub>3</sub> transfer for other N<sub>r</sub> instruments (Crilley et al., 2023). In

696 contrast, during the emptying process, desorption occurs over previously exposed surfaces ending  
 697 at the instrument, reducing the extent of surface effects. The surface interactions for these gases  
 698 are minimized in the modified chambers to facilitate more time-efficient measurements of surface  
 699 exchange. However, they necessitate the use of the  $\lambda$  term when deployed in the measurement-  
 700 reference configuration for those  $N_r$  species which experience partial transmission, such as  $NH_3$ .  
 701 The  $\lambda$  term is required to obtain accurate values, as the enclosed flux measurement surface should  
 702 be perturbed for the least amount of time possible when making field measurements, and the  
 703 chambers cannot be closed for several hours to allow surface-active gases to passivate the lines.  
 704 One potential option to improve the system performance further for  $NH_3$  could be to heat the gas  
 705 transfer lines between the chambers and gas analyzers. In addition, minimizing the potential for  
 706 transformations reduces the frequency required for in-field characterization of these processes  
 707 through positive and negative gas delivery controls. For  $NO_2$ , specifically, we sought to quantify  
 708 this as a function of modifying components of our chambers, as  $NO_2$  is the most reactive gas in  
 709 our suite (Section S5).

710 **Table 1.** Summary of time responses of addition and removal of GHGs and  $N_r$  gases to and from  
 711 a chamber at  $2 \text{ L min}^{-1}$  in unmodified and modified configurations. Time responses here  
 712 correspond to a theoretical fill or empty e-folding time response of 36 minutes. Where analytes  
 713 were observed to undergo surface interactions, a double exponential fit was used to characterize  
 714 them, with the first time constant representing the known gas exchange rate being fixed at 36  
 715 minutes, as this was observed for the non-reactive gases, and the second time constant is reported  
 716 (\*) alongside an assessment of the relative magnitude of surface interactions, through the D-value  
 717 (%) (Crilley et al., 2023; Ellis et al., 2010; Moravek et al., 2019). Variability shown is one standard  
 718 deviation of the mean from replicate experiments (n=3).

Gas species	Direction	Unmodified (min)	Modified (min)	Improvement (min)	D (%)
NO	Fill	$38.8 \pm 0.7$	$37.8 \pm 0.6$	$1.0 \pm 0.7$	-
	Empty	$37.9 \pm 1.5$	$36.0 \pm 0.6$	$1.9 \pm 2.1$	-
$NO_2$	Fill*	$18.9 \pm 0.6$	$18.0 \pm 3.1$	$0.9 \pm 3.2$	$94 \pm 18$
	Empty*	$21.2 \pm 1.2$	$20.4 \pm 1.4$	$0.8 \pm 1.9$	$78 \pm 21$
HONO	Fill*	$21.9 \pm 1.1$	$19.3 \pm 1.2$	$2.6 \pm 1.6$	$74 \pm 10$
	Empty*	$23.2 \pm 1.4$	$21.2 \pm 0.9$	$2.0 \pm 1.7$	$71 \pm 9$

NH <sub>3</sub>	Fill*	69.6 ± 0.4	68.2 ± 0.5	1.4 ± 0.6	89 ± 3
	Empty*	76.9 ± 0.8	75.0 ± 4.6	1.9 ± 4.7	23 ± 4
CO <sub>2</sub>	Fill	37.9 ± 1.5	37.0 ± 1.8	0.9 ± 2.3	-
	Empty	38.0 ± 1.8	37.0 ± 1.2	1.0 ± 2.2	-
CH <sub>4</sub>	Fill	37.9 ± 1.1	36.8 ± 1.2	1.1 ± 1.6	-
	Empty	39.1 ± 1.7	37.2 ± 1.2	1.9 ± 2.1	-
N <sub>2</sub> O	Fill	38.6 ± 2.1	36.7 ± 1.2	1.9 ± 2.4	-
	Empty	39.7 ± 1.3	37.7 ± 1.6	2.0 ± 2.1	-

719

### 720 3.2 Multiplexer modification impacts on gas transfer

721 The multiplexer (eosMX; Eosense Inc.) allows ~~us to operation of~~ up to twelve dynamic chambers  
722 simultaneously with a suite of gas analyzers ~~such as the Picarro G2509, NO<sub>x</sub>, and O<sub>3</sub> analyzers.~~  
723 However, ~~the commercial multiplexer it~~ is constructed with stainless steel (SS) valves and fittings  
724 that would be expected to facilitate strong interactions ~~and/or~~ losses of target gases in the N<sub>r</sub> analyte  
725 suite. Valves and fittings made of SS have a higher tendency to chemically interact and/or adsorb  
726 reactive gases compared to fluoropolymer replacements. To address this uncertainty, the gas  
727 transfer efficiency as a percentage loss in the multiplexer versus a bypass line was evaluated  
728 specifically for NH<sub>3</sub> and NO<sub>2</sub>, alongside standard GHGs as they passed through fittings and gas  
729 handling solenoid valves made of SS or PFA and PTFE replacement parts.

730 The loss fractions were modest and measurable when ~~connected to the analyzers~~ using minimal  
731 lengths of PFA tubing (~50 cm) ~~instead of the standard 15 m gas transfer lines.~~ The most  
732 substantial loss ~~of gases~~ was observed on SS, ~~as expected in which the highest reactivity and/or~~  
733 ~~adsorption were expected~~ due to its known tendencies ~~to interact with reactive gases~~ (Vaaitinen  
734 et al., 2014). When the GHGs were delivered ~~after dry zero air into the SS setup~~ for 30 minutes,  
735 typical of a chamber closure period ~~in the field~~, their losses ranged from 10% for N<sub>2</sub>O to 19% for  
736 H<sub>2</sub>O. Meanwhile, NO<sub>2</sub> exhibited 17% loss on the SS surfaces, and the greatest effect was seen for  
737 NH<sub>3</sub> with a loss of 38% (Figure S6). In contrast, losses on the chemically inert and hydrophobic  
738 surface of the PFA fittings and PTFE valve were negligible (<1%) for most gases, except for NH<sub>3</sub>,  
739 which still exhibited a measurable loss of 11%. Other reports have also shown up to 15% loss of

740 NH<sub>3</sub> at atmospheric pressure on PTFE and PFA surfaces (Ellis et al., 2010; Shah et al., 2006;  
741 Vaittinen et al., 2014). While it is expected that the SS would eventually passivate and improve  
742 the transmission of the GHGs in a standard recirculation approach, this is not likely to be the case  
743 for destructively analyzed N<sub>r</sub> and even more so if it facilitates a chemical transformation. Overall,  
744 we found that replacing the multiplexer SS valves and fittings with PFA fittings and PTFE valves  
745 provided a 9–27% reduction in surface losses of N<sub>r</sub> compounds and GHGs. We strongly  
746 recommend the use of PTFE and/or PFA materials over SS for more accurate measurement of N<sub>r</sub>  
747 species when interfacing the dual chamber setup with the destructive N<sub>r</sub> analyzers needed for field  
748 flux measurements, whether using a custom setup or the commercially available multiplexer.

749

750

### 751 **3.3 Minimizing NO<sub>2</sub> losses and determining controlling variables**

752 In addition to the rate of transfer of N<sub>r</sub> gases, chamber modifications are necessary to prevent  
753 reactive losses. ~~The lost fraction of NO<sub>2</sub> ( $f_{\text{NO}_2}$ ) to chamber surfaces was quantified by the addition  
754 of known mixing ratios of calibration gas to the chamber in both the unmodified and modified  
755 configurations under atmospherically relevant humidities.~~ These experiments enabled us to  
756 determine the magnitude of NO<sub>2</sub> lost to the chamber and gas transfer tubing surfaces due to  
757 chemical and/or physical transformations, as well as demonstrate the effectiveness of the different  
758 chamber modifications in minimizing these losses. For NO<sub>2</sub>, a probable chemical transformation  
759 pathway is its heterogeneous conversion to HONO (R3), which is favourable under  
760 atmospherically relevant humidities, and the resulting water-adsorbed surfaces expected to exist  
761 throughout the chamber and sampling lines.

#### 762 **3.3.1. Chamber modification impacts on NO<sub>2</sub> losses**

763 ~~The average fractional loss ( $f_{\text{NO}_2}$ ) for an unmodified chamber and one modified with the PFA film  
764 and PTFE bulkhead fittings was measured under conditions of 83% RH and 5 ppb of NO<sub>2</sub>.~~  
765 Substantial reduction in NO<sub>2</sub> loss fraction ( $f_{\text{NO}_2}$ ) and transformation was observed from the  
766 implemented PFA and PTFE modifications under conditions of 83% RH and 5 ppb of NO<sub>2</sub>. In the  
767 original unmodified configuration of the chamber,  $f_{\text{NO}_2}$  was  $0.36 \pm 0.02$  (Figure S7). ~~When the  
768 interior chamber surfaces were covered with a film of PFA, it which~~ was reduced to  $0.22 \pm 0.03$

769 [with the PFA film](#), a relative decrease of 18%. This is consistent with the acrylic chamber surfaces  
770 and fasteners to the chamber frame, facilitating physical and/or chemical loss of NO<sub>2</sub>.

771 The film of PFA, along with other fluoropolymers, is known to have excellent chemical resistance  
772 and low reactivity towards a range of chemicals, including NO<sub>2</sub> (Ebnesajjad, 2005; Graham et al.,  
773 1997). In addition, the superhydrophobic nature of these materials prevents the accumulation of  
774 water on the surfaces, which is known to facilitate atmospheric surface reactions of NO<sub>2</sub>  
775 (Finlayson-Pitts et al., 2002; Jenkin et al., 1988; Stutz et al., 2002) and create analytical bias in the  
776 measurement of trace gases like HONO, especially when instrument gas sampling inlets do not  
777 take this into account (Crilley et al., 2019; Von Der Heyden et al., 2022).

778 The replacement of the brass-lined push-to-connect bulkhead fittings with PTFE led to a similar  
779 decrease in  $f_{\text{NO}_2}$ , which was reduced by 17% to a final value of less than  $0.05 \pm 0.02$  (Figure S7).  
780 These fitting surfaces act as the largest surface-driven NO<sub>2</sub> loss despite their surface area being  
781 very small compared to that of the entire chamber configuration and with a very small contact time  
782 against the gas sample (0.012 s per fitting at a flow rate of 2 L min<sup>-1</sup>).

783 The loss of NO<sub>2</sub> in the commercially available system is challenging to attribute solely to the  
784 heterogeneous hydrolysis reaction. During the characterization experiments, the conditions inside  
785 the chamber were matched to those reported by previous lab studies, which have shown that high  
786 RH, presence of NO<sub>2</sub>, and surface adsorbed water on surfaces favour this loss mechanism (Jenkin  
787 et al., 1988; Stutz et al., 2004). The reaction is known to occur on surfaces such as Pyrex (Jenkin  
788 et al., 1988) and borosilicate glass (Finlayson-Pitts et al., 2002), but no prior studies to date, nor  
789 this study, have demonstrated metallic surfaces as facilitating this mechanism.

790 ~~Overall, chamber modifications were effective in minimizing NO<sub>2</sub> losses.~~ The inert PTFE fittings  
791 dramatically minimized transformations, while PFA film lining the inner chamber surfaces was  
792 also effective, but less so. Our results indicate that water-adsorbed and metallic surfaces, such as  
793 brass, facilitate substantial loss and/or transformations of NO<sub>2</sub>. Further investigation is required to  
794 confirm the mechanism(s) at play and is beyond the scope of this work.

795

### 796 **3.3.2. RH-facilitated NO<sub>2</sub> loss as a function of concentration**

797 A complete characterization of  $f_{\text{NO}_2}$  and the amount of HONO in the fully modified chamber was  
798 determined under a range of environmentally relevant RHs and  $\text{NO}_2$  concentrations. We found that  
799 the absolute and fractional  $\text{NO}_2$  losses were highest under the highest RH conditions (85%; Table  
800 2). However, the  $f_{\text{NO}_2}$  does not appear to follow a concentration-dependent trend across the  
801 additions made at lower RHs, with at most 0.4 ppbv  $\text{NO}_2$  lost across the remainder of the tests, a  
802 value which is equivalent to the LOD of the  $\text{NO}_x$  analyzer used. This would generate 0.2 ppbv of  
803 HONO according to the disproportionation of the hydrolysis mechanism, which is well below the  
804 analyzer detection limits. ~~Overall, the modifications successfully reduced~~  $\text{NO}_2$  losses below 10%  
805 across all environmentally relevant conditions the chambers are expected to encounter, with our  
806 findings here suggesting that the mass lost is nearly constant and independent of  $\text{NO}_2$  mixing ratio  
807 at RHs below 85%, while being marginally higher at and above this value.

808 Quantifying  $f_{\text{NO}_2}$  and the amount of HONO made in the chamber is required for the correction of  
809 field datasets. ~~Quantifying  $\text{NO}_2$  loss to the chamber surfaces allows us to prevent bias when~~  
810 ~~reporting  $\text{NO}_2$  exchange processes on environmental surfaces, as the magnitude of the losses~~  
811 ~~attributed to chamber surfaces versus the surface studied can be corrected.~~ The dual chamber  
812 system, via the reference chamber RC, can also quantify any changes in these processes over time  
813 if standard ~~concentration~~ additions to the headspace are conducted. Consequently, important  
814 parameters such as  $\text{NO}_2$  deposition fluxes on surfaces can be better estimated (Pape et al., 2009).

815 Since the inferred HONO mixing ratios from the chamber surfaces across various environmental  
816 RHs are nearly invariant at  $0.2 \pm 0.1$  ppbv, it is simple to background correct any observational  
817 datasets by subtracting this amount from the total HONO measured in the chamber. ~~if a more~~  
818 ~~sensitive analyzer is used during field or lab experiments.~~ In addition, as the  $\text{NO}_2$  values expected  
819 in most atmospheric gas samples during field measurements are well into the ppbv range ( $>3.3$   
820 ppbv per 30-minute flux measurement for a  $1.40.08 \mu\text{g N m}^{-2} \text{ min}^{-1}$  emission), the corrections  
821 would be easy to implement in post-processing of datasets and have minimal impact on the  
822 technical aspects of the analytical determinations.

823 **Table 2.** Characterization of  $\text{NO}_2$  lost in the modified chamber under environmentally relevant  
824 ranges of  $\text{NO}_2$  and RH. The loss fraction ( $f_{\text{NO}_2}$ ) and HONO produced in the chamber were  
825 quantified ~~at each of three  $\text{NO}_2$  mixing ratios (5, 7, 10 ppbv) delivered at varying relative~~  
826 ~~humidities (RH) to the fully modified chamber.~~ Variability ( $\pm$ ) provided is one standard deviation  
827 of the mean from replicate experiments ( $n=3$ ).

RH (%)	NO <sub>2</sub> added (ppbv)	NO <sub>2</sub> lost (ppbv)	$f_{NO_2}$	HONO produced (ppbv)
85	5	0.50 ± 0.01	0.1 ± 0.02	
85	7	0.70 ± 0.04	0.1 ± 0.02	
85	10	0.50 ± 0.07	0.05 ± 0.01	
65	5	0.30 ± 0.10	0.06 ± 0.02	
65	7	0.40 ± 0.09	0.06 ± 0.02	<1.1 <sup>a</sup>
65	10	0.30 ± 0.08	0.03 ± 0.01	
45	5	0.30 ± 0.05	0.05 ± 0.01	
45	7	0.40 ± 0.04	0.06 ± 0.00	
45	10	0.30 ± 0.08	0.03 ± 0.01	

<sup>a</sup> – below instrument detection limit of 1.1 ppbv determined as S/N=3 while sampling zero air

It should be noted that the amount of HONO in the chamber was below the LOD of the NO<sub>x</sub> analyzer for HONO (1.071 ppbv), meaning that the upper limit of HONO inferred may perhaps, in fact, be negligible. Therefore, future experiments that wish to detect small N<sub>r</sub> fluxes accurately will need to focus on reproducing these experiments with a higher performance instrument, such as a time-of-flight chemical ionization mass spectrometer (ToF-MS) or long-path absorption photometer (LOPAP), which have lower detection limits (Crilley et al., 2019; Lee et al., 2014; Neuman et al., 2016; Reed et al., 2016).

### 3.23.3 Characterization Loss of O<sub>3</sub> loss in the chamber modified with and without PFA fluoropolymer modifications

Ozone loss was observed in both modified and unmodified chamber configurations, with 18% lost to clean 15 m PFA lines alone when transferring 30 ppbv. The unmodified chambers lost 45% across ~~all~~ delivered mixing ratios spanning (150-250 ppbv), which are higher than those typically expected in ambient air. This was reduced to 35% when the fluoropolymer modifications were implemented ~~with new clean PTFE fittings and PFA film~~. When the PFA film was aged by 15 days of ambient sampling and 2 years of exposure to lab air, the losses were substantially exacerbated, reaching 80%. Such outcomes are expected and can be attributed to several factors discussed in detail in Section S5, primarily involving surface reactions with built-up films of

864 deposited organics, adsorption, and material interactions (Burkholder et al., 2015.) [Ebnesajjad,](#)  
865 [2017, George et al., 2015, Plake et al., 2015](#)). [We recommend regular replacement of the PFA film](#)  
866 [as part of the N<sub>r</sub> system maintenance, coupled with quality control procedures to characterize](#)  
867 [material performance for target gases.](#)

868 ~~Ozone is known to undergo heterogeneous surface reactions, particularly on materials like glass,~~  
869 ~~metals, or polymers (Plake et al., 2015). These surfaces often contain reactive sites such as~~  
870 ~~hydroxyl groups or adsorbed water molecules that can catalyze the decomposition of O<sub>3</sub> into~~  
871 ~~molecular oxygen (O<sub>2</sub>) and other byproducts (George et al., 2015). Adsorption of O<sub>3</sub> on chamber~~  
872 ~~surfaces is another potential pathway for loss. When O<sub>3</sub> molecules interact with surfaces, they may~~  
873 ~~undergo a reaction if the materials are not inert. The use of PFA here was intended to benefit the~~  
874 ~~transfer of O<sub>3</sub>, yet despite its high chemical inertness, low reactivity, and resistance to the uptake~~  
875 ~~of many chemicals from gas samples, the subsequent O<sub>3</sub>-driven surface reactions are still~~  
876 ~~substantial (Ebnesajjad, 2017). This is likely because, over time, reactive substances from sampled~~  
877 ~~air accumulated through adsorption on the PFA film and/or its defect sites, thereby reducing its~~  
878 ~~effectiveness. The same issue is well known for PFA tubing used in standard air quality monitoring~~  
879 ~~and is a common maintenance need for O<sub>3</sub>-analyzer inlets. Further, physical wear or surface aging~~  
880 ~~might alter the material surface through product formation or exposure of new reaction sites, which~~  
881 ~~makes it less resistant to further reaction with O<sub>3</sub>, and therefore increases the rate of loss over time.~~  
882 ~~This is the case in our aged PFA film results, which emphasizes the importance of regular~~  
883 ~~replacement of the film as part of the N<sub>r</sub> system maintenance and that quality control procedures~~  
884 ~~characterizing the chamber material performance with respect to O<sub>3</sub> will provide the highest~~  
885 ~~accuracy of experimental results.~~

886

### 887 **3.3 Proof-of-concept reactive nitrogen fluxes using soil samples in the lab**

888 Proof-of-concept flux measurements were performed using the modified dynamic chamber system  
889 to demonstrate that emissions of N<sub>r</sub> gases from agricultural soil samples can be measured under  
890 controlled conditions, similar to many prior reports using custom-built soil chambers (Almand-  
891 Hunter et al., 2015; Pape et al., 2009; Tang et al., 2019, 2020).

#### 892 **3.3.1. Fluxes of NO, NO<sub>2</sub>, and HONO from agricultural soil samples**

893 Emission fluxes were measured from two pooled and two individual soil samples collected from  
894 a single agricultural field (Table 3; Section S6). The average and integrated fluxes of N<sub>2</sub>O, NH<sub>3</sub>,  
895 NO, NO<sub>2</sub>, and HONO were assessed under controlled, environmentally realistic (65 % RH), drying  
896 conditions of zero air modified to be 65% RH delivered at 3.6 L min<sup>-1</sup> to the chamber headspace  
897 where the soil was contained. The soil started the drying process from a VWC of approximately  
898 25% with the flows held constant for around 4 days or until the VWC reached 15% (Table 3).

899 As the soils dried, NO and NO<sub>2</sub> emissions increased, with NO fluxes highest across all replicates  
900 and reaching up to 2.50 µg N m<sup>-2</sup> hr<sup>-1</sup>. This trend is consistent with the prior work of other  
901 researchers, showing peak NO emission potentials when VWC drops below 25% during soil  
902 drying, which is when microbial nitrification and denitrification processes are suggested to become  
903 more active (Bao et al., 2022; Oswald et al., 2013). The soil VWC at which these maxima occur  
904 can vary depending on soil type, texture, and the microbial diversity therein (Ludwig et al., 2001;  
905 Schindlbacher et al., 2009). The plot-level replicates from our field had a higher integrated NO  
906 flux (i.e., > 2600 µg N m<sup>-2</sup>), compared to the pooled replicates (<1000 µg N m<sup>-2</sup>), likely indicating  
907 real differences in preserved microbial hotspots, intact plot-level soil aggregates, true spatial  
908 variability, and plot-specific N availability (Table 3). Soil texture and aggregate size, for example,  
909 play an important role in building the porous structure of soil, which has implications for the  
910 release of gases (Mangalassery et al., 2013). Soil aggregates, therefore, govern the release of  
911 gaseous N<sub>r</sub> analytes like NO based on the aerobic or anaerobic state of the soil. Here, our low level  
912 of soil manipulation (i.e. not ground, no sieving) will drive some of the variability by preserving  
913 these features, which exist across and within real soil systems (Lipiec et al., 2007). Individual plot  
914 samples also retain plot-specific microbial communities when working with intact soil, whereas  
915 soil grinding can temporarily inhibit microbial activity. While we tried to minimize soil handling  
916 and processing extremes in these experiments, a measure of homogeneity was also pursued, and  
917 fully intact soil cores were not assessed.

918 Integrated NO<sub>2</sub> fluxes showed the same trend, with more sample-to-sample variability as one  
919 pooled replicate (R2) produced over three times the emissions of R1, despite both experiments  
920 being conducted across identical moisture content ranges (Table 3). Given the limited studies  
921 directly measuring NO<sub>2</sub>, such as Purchase et al. (2023), this variability is difficult to interpret and

922 highlights the need for more assessments of its production pathways and controls, which our  
923 developed chambers show promise for.

924 The observed average HONO fluxes remained low across all of the samples, ranging from 0.05 to  
925  $0.25 \mu\text{g N m}^{-2} \text{ hr}^{-1}$  (Table 3). These values are lower by ~~up to more than~~ an order of magnitude  
926 compared to those reported in other controlled laboratory studies, where HONO fluxes ~~were~~  
927 ~~around-exceed~~  $900 \mu\text{g N m}^{-2} \text{ shr}^{-1}$  (Oswald et al., 2013; Su et al., 2011; Wang et al., 2021). These  
928 discrepancies are concerning, given recent emphasis from the scientific community on the  
929 atmospheric impacts of soil-derived HONO on air quality. Here, the results from our agricultural  
930 soil samples may reflect the differences in our methodology, such as the soil handling and  
931 preparation steps prior to and during experiments. Many prior reports prepare their samples in  
932 ways that strongly deviate from real-world conditions (e.g. initial soil drying temperatures above  
933 those occurring under ambient conditions, extreme storage conditions, grinding, sieving, use of  
934 dry zero air to flush chambers, etc.). Further drivers of variability within the category of heavily  
935 altered soil samples from the literature include pH,  $\text{NO}_2^-$  availability, and  $\text{NH}_4^+$  or  $\text{NO}_3^-$  content,  
936 all of which are known to influence biotic and abiotic HONO formation pathways (Wu et al.,  
937 2019).

938 Our HONO fluxes from the agricultural soil samples studied here are consistent with field  
939 observations under ambient conditions, where average emissions have been reported to largely  
940 remain below  ~~$2 \text{ ng N m}^{-2} \text{ s}^{-1}$~~  or  $7.2 \mu\text{g N m}^{-2} \text{ hr}^{-1}$  (Tang et al., 2019; Xue et al., 2024). This does  
941 suggest that greater care in sample preparation, and likely also a widely agreed-upon standard  
942 procedure, is needed to study soil HONO emissions relevant to atmospheric models.

943 The integrated HONO fluxes for the pooled replicates yielded  $185 \mu\text{g N m}^{-2}$  and  $146 \mu\text{g N m}^{-2}$ ,  
944 respectively. From the individual sample replicates, which were slightly wetter than the pooled,  
945 the integrated HONO fluxes were 690 and  $739 \mu\text{g N m}^{-2}$ , which was unexpected because the  
946 overall moisture regime accessed by the pooled soil experiments was not lower than those from  
947 the plot samples. The drier soils would have been expected to yield greater integrated HONO  
948 emissions (Oswald et al., 2013), yet this was not the case. Additional replicates and experimental  
949 controls, while beyond the scope of this study, would allow further attribution of the controls over  
950 the observed HONO variability.

951 **Table 3.** Average and integrated fluxes of NO, NO<sub>2</sub>, and HONO (in  $\mu\text{g N m}^{-2} \text{hr}^{-1}$  and  $\mu\text{g N m}^{-2}$ , respectively) from agricultural soil  
 952 samples across two soil VWC ranges. Both the average and integrated fluxes were calculated over a constant period within the noted  
 953 range of soil VWC. Values are reported as mean  $\pm$  standard error.

Soil sample	Soil VWC Range (%)	Duration (hr)	Avg Flux ( $\mu\text{g N m}^{-2} \text{hr}^{-1}$ )			Integrated flux ( $\mu\text{g N m}^{-2}$ )		
			NO	NO <sub>2</sub>	HONO	NO	NO <sub>2</sub>	HONO
Pooled R1	16-22	125	1.0 $\pm$ 0.04	0.05 $\pm$ 0.02	0.06 $\pm$ 0.02	2400 $\pm$ 120	160 $\pm$ 50	190 $\pm$ 50
Pooled R2	16-22	125	1.0 $\pm$ 0.04	0.20 $\pm$ 0.02	0.05 $\pm$ 0.01	2500 $\pm$ 130	500 $\pm$ 60	150 $\pm$ 40
Plot R1	22-27	65	1.2 $\pm$ 0.05	0.40 $\pm$ 0.03	0.20 $\pm$ 0.02	3400 $\pm$ 150	1300 $\pm$ 90	690 $\pm$ 70
Plot R2	22-27	65	1.0 $\pm$ 0.03	0.50 $\pm$ 0.04	0.30 $\pm$ 0.02	2600 $\pm$ 100	1600 $\pm$ 100	740 $\pm$ 50

954

955 These findings demonstrate the utility of the modified custom-built dynamic chambers for  
956 accurately capturing  $N_r$  fluxes under controlled laboratory settings, but they also highlight the need  
957 for more such systems to be implemented across the scientific community to better consider both  
958 biogeochemical soil properties and environmental context when interpreting the impacts of  $N_r$   
959 fluxes obtained in the lab and scaling them to real soils. There seems to be potential for skewing  
960 the atmospheric impacts as a result, in particular for HONO, as the standard approaches have been  
961 designed to replicate NO fluxes (Behrendt et al., 2014). Most global models do not consider the  
962 effect of soil HONO on air quality through  $O_3$  production and oxidation chemistry (Ha et al., 2023).  
963 Several modelling studies like Ha et al. (2023) and Tian et al. (2024) have incorporated the order  
964 of magnitude or higher HONO fluxes reported from lab studies, like those by Su et al. (2011),  
965 Wang et al. (2021), Oswald et al. (2013), and Meusel et al. (2018). They estimated significant  
966 HONO production with maximum flux potentials of  $830 \mu\text{g N m}^{-2} \text{hr}^{-1}$ ,  $95 \mu\text{g N m}^{-2} \text{hr}^{-1}$ ,  $70 \mu\text{g N}$   
967  $\text{m}^{-2} \text{hr}^{-1}$ ,  $55 \mu\text{g N m}^{-2} \text{hr}^{-1}$ , respectively. In contrast, the field observations that do exist suggest that  
968 real HONO fluxes are much smaller at  $2\text{-}17.5 \mu\text{g N m}^{-2} \text{hr}^{-1}$  (Song et al., 2023; Tang et al., 2019).  
969 Similarly, Wu et al. (2022) has used the regional WRF-Chem model to explore the impact of soil  
970 HONO emissions on the concentrations of atmospheric HONO, OH, and  $O_3$ .

971 Agricultural soil HONO emissions have been suggested to significantly contribute to OH radical  
972 production, accounting for approximately 10% to 60% of total OH formation in rural areas before  
973 noon (Oswald et al., 2013; Su et al., 2011), which often exceed the contributions from ozone  
974 photolysis. Additionally, high HONO emissions from agricultural soils have been reported to  
975 increase local  $O_3$  concentrations by  $\sim 0.5 - 1.0$  ppb in low- $\text{NO}_x$  rural environments where VOCs  
976 are not limiting (Zhang et al., 2021), with even greater impacts observed during fertilization  
977 periods (Wu et al., 2022). Modelling studies, using GEOS-Chem and CMAQ, for example, claim  
978 that incorporating soil HONO emissions improves the agreement between observed and simulated  
979  $O_3$  levels, particularly during the morning (Zhang et al., 2021).

980 Only a few studies have conducted estimates of soil  $\text{NO}_x$  emissions under reasonable conditions,  
981 like Bao et al. (2022) and Wu et al. (2022), where the researchers tried to mimic field conditions.  
982 Even in such an area that has been long studied, large uncertainty still exists around soil sources  
983 of  $\text{NO}_x$ , particularly for agricultural activities where uncertainty is still at least  $\pm 30\%$  due to  
984 limitations in lab characterizations and field experiments (Gong et al., 2025). It is not surprising,

985 then, that a similar issue exists for the very recent work on HONO soil emissions. The NO<sub>x</sub>  
986 uncertainty range results from intricate soil biogeochemical processes and varies with crop types,  
987 soil texture, fertilizer types and application rate (Gong et al., 2025). This longstanding and  
988 established difficulty in predicting soil NO<sub>x</sub> for use in global chemical models means that doing  
989 so for HONO without careful ground truthing of real-world emissions could lead to substantial  
990 inflation of the impacts on atmospheric chemistry and air quality. Care should be taken in using  
991 HONO emissions from lab studies in global models, as it seems they pose a risk of overestimating  
992 their atmospheric impacts until a more representative experimental design can be obtained with  
993 chamber systems like the one used here.

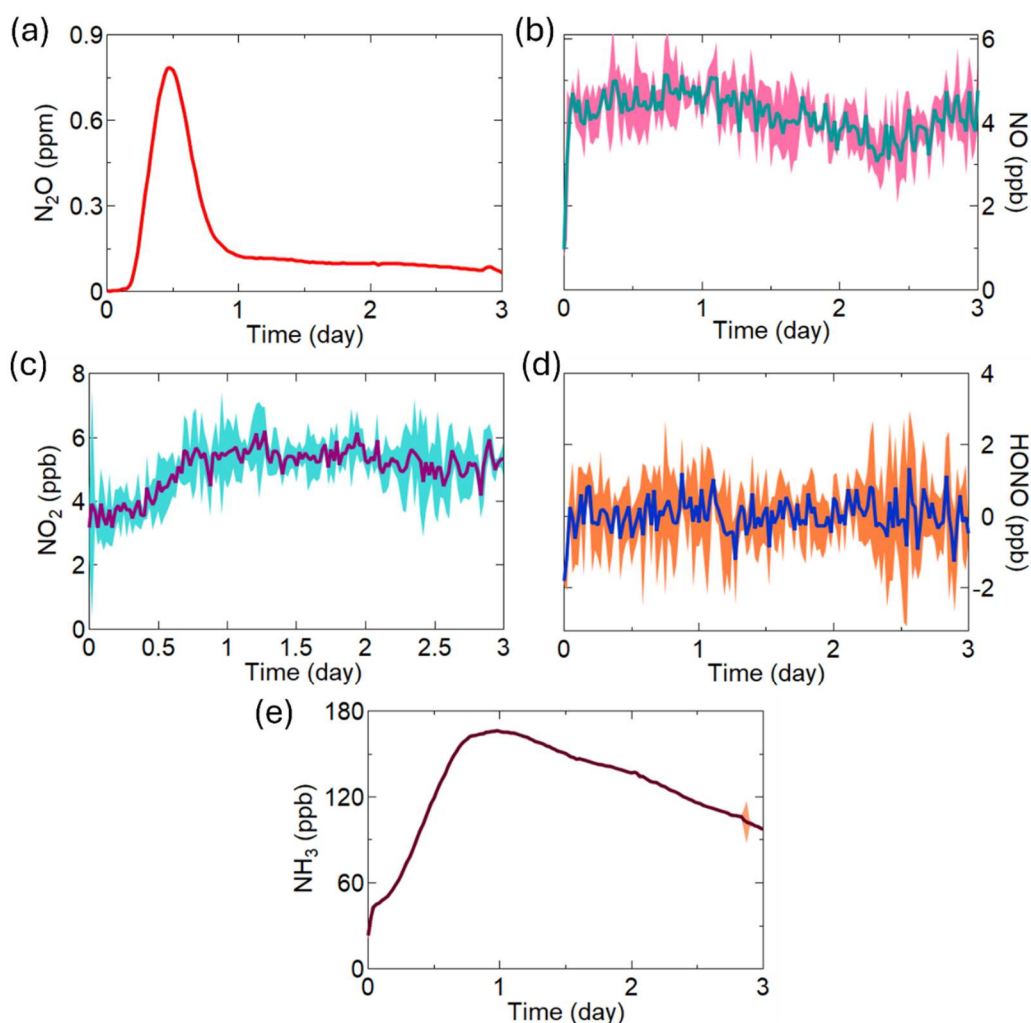
994

### 995 **3.3.2 Fluxes of NH<sub>3</sub>, N<sub>2</sub>O, NO, NO<sub>2</sub>, and HONO/N<sub>r</sub> from fertilized agricultural soil samples**

996 Agricultural soils amended with chemical fertilizers are expected to be hotspots for NH<sub>3</sub>, N<sub>2</sub>O,  
997 HONO, and NO<sub>x</sub> emissions. Here, we demonstrate the use of our developed chambers to measure  
998 these analytes under controlled lab conditions using our lightly processed pooled soil samples and  
999 four fertilizers: urea, ammonium nitrate (AN), ammonium bicarbonate (ABC), and ammonium  
1000 carbonate (AC) with the temperature maintained at 23 °C, VWC between 25–29%, and headspace  
1001 RH held at 65% for three days, simulating realistic atmospheric and environmental N<sub>r</sub> flux  
1002 exchange conditions following farm field fertilization (Figure 3). In all experiments conducted  
1003 with the Picarro G2509, the mixing ratio of N<sub>2</sub>O began to rise approximately four hours after the  
1004 experiment started. In the example shown for urea, it peaked at 0.79 ppm after 12 hours, which  
1005 was followed by a gradual decline (Figure 3a). This pattern likely reflects the incubation period of  
1006 nitrifying and denitrifying bacteria that leads to the subsequent release of gases like HONO, as  
1007 depicted in Wang et al. (2021), and N<sub>2</sub>O in Liu et al. (2022). In contrast, the mixing ratio of NO  
1008 remained relatively constant throughout the three days (Figure 3b), with NO<sub>2</sub> increasing as the  
1009 N<sub>2</sub>O emissions decreased, while NO was emitted constantly throughout (Figure 3c). In this  
1010 example experiment, no measurable emissions of HONO were detected despite the substantial  
1011 presence of urea and evidence of active microbial nitrification and denitrification from the other  
1012 emitted gases (Figure 3d). Lastly, the urea application example in Figure 3e shows the expected  
1013 significant NH<sub>3</sub> emissions, with the integrated amount reaching 22% of the applied N over the

1014 three-day incubation period. These findings are consistent with our existing knowledge that  $\text{NH}_3$   
1015 volatilization as an N loss mechanism dominates early  $\text{N}_r$  losses from fertilizers.

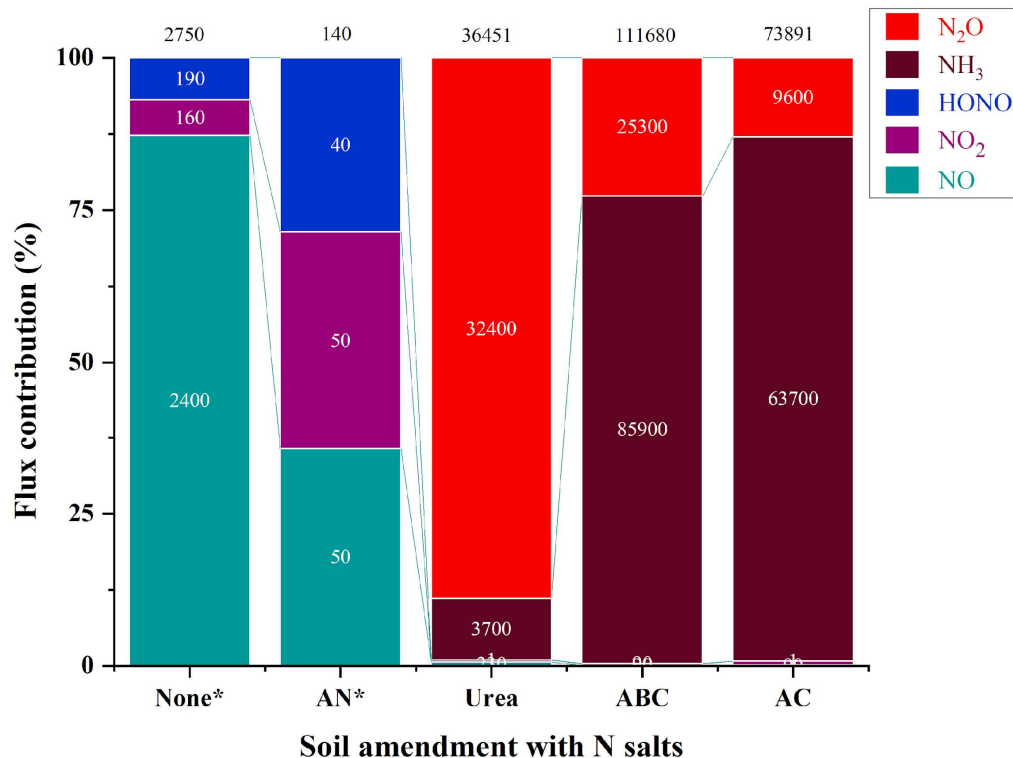
1016 Volatilization of  $\text{NH}_3$  is well-characterized as a major pathway for N loss from fertilizers (Behera  
1017 et al., 2013; Govoni Brondi et al., 2024; Liu et al., 2020; Moravek et al., 2019; Pan et al., 2016,  
1018 2022; Paulot et al., 2014). Besides agronomic concerns due to N loss and reduced fertilizer  
1019 efficiency related to  $\text{NH}_3$  emissions from fertilized soil (Anas et al., 2020), it is a key precursor to  
1020 secondary inorganic aerosols in the atmosphere with impacts on respiratory and ecosystem health,  
1021 visibility, and climate (Dennis et al., 2010; Edwards et al., 2024; Fowler et al., 2013; González  
1022 Ortiz et al., 2020; Jang et al., 2025; Seinfeld and Pandis, 2006).



1023  
1024 **Figure 3.** Soil emissions of the comprehensive suite gases from soil treated with urea, including  
1025 (a)  $\text{N}_2\text{O}$ ; (b)  $\text{NO}$ ; (c)  $\text{NO}_2$ ; (d)  $\text{HONO}$ ; and (e)  $\text{NH}_3$ . The  $\text{NO}$ ,  $\text{NO}_2$  and  $\text{HONO}$  emissions were  
1026 measured at 1-minute resolution and averaged to 30 min. The standard deviation ( $\pm 2\sigma$ ) around  
1027 these averages is shaded with a lighter version of the main trace colour. Similarly, the 2-second

1028 resolution of N<sub>2</sub>O and NH<sub>3</sub> measurements were also averaged to 30-minute intervals with ±1σ  
1029 provided in shading.  
1030  
1031 Emissions of NH<sub>3</sub> and N<sub>2</sub>O were observed to be far greater in terms of integrated amounts from  
1032 the fertilized samples (Figure 4). Integrated flux for NH<sub>3</sub> produced by soil treated with ABC  
1033 accounted for 77% of total N<sub>r</sub> flux (i.e., 85\_900 μg N m<sup>-2</sup>), followed by AC, which was 63\_700 μg  
1034 N m<sup>-2</sup>. This is not surprising, since the use of chemical fertilizers increases the concentration of  
1035 NH<sub>4</sub><sup>+</sup> in the soil that can deprotonate to emit neutral NH<sub>3</sub> and, in the presence of ammonia-  
1036 oxidizing microorganisms, increases the production of N<sub>2</sub>O (Luo et al., 2025). Nitrous oxide  
1037 released from the addition of urea accounted for the highest integrated flux of 32\_400 μg N m<sup>-2</sup>  
1038 observed, representing 89% of total N<sub>r</sub> released, followed by 25\_300 μg N m<sup>-2</sup> and 9\_600 μg N m<sup>-2</sup>  
1039 for ABC and AC, respectively. One way that has been proposed to reduce these large N<sub>2</sub>O  
1040 emissions from inorganic fertilizers is to change the application form to organic fertilizer, as the  
1041 NH<sub>4</sub><sup>+</sup> in soil is ~~released~~produced more slowly (Luo et al., 2025). Due to a limited duration of  
1042 access to the G2509 to conduct this work, we were unable to measure the N<sub>2</sub>O and NH<sub>3</sub> emitted  
1043 from unamended soils or those treated with AN. Regardless, based on the obtained data, the values  
1044 found here are similar to those observed under real environmental conditions (Figure 4). For our  
1045 sample without N amendment, the integrated flux of NO was the largest (2\_400 μg N m<sup>-2</sup>; 87% of  
1046 total NO<sub>y</sub>), followed by comparable levels of NO<sub>2</sub> (160 μg N m<sup>-2</sup>) and HONO (190 μg N m<sup>-2</sup>). The  
1047 fluxes of NO<sub>x</sub> and HONO were below 1 μg N m<sup>-2</sup> hr<sup>-1</sup> for all the nutrient addition treatments,  
1048 suggesting a similar trend as those observed under field conditions and from our lab results with  
1049 unamended soils (Figure 4, Table S1, Section S6). The exact mechanism behind the HONO  
1050 release, being due to nitrification and/or denitrification, cannot be definitively assigned based on  
1051 flux data alone, and many ~~abiotic~~ factors ~~like soil moisture, temperature, and presence of specific~~  
1052 ~~microbial communities~~ drive these emissions. There are discrepancies still observed between  
1053 HONO flux measured from the treated soil samples in the laboratory and similar measurements in  
1054 literature (Oswald et al., 2013; Su et al., 2011), some of which report fluxes of up to ~3\_600 μg N  
1055 m<sup>-2</sup> hr<sup>-1</sup> and are likely overestimating the soil N<sub>r</sub> fluxes found in the real world. In contrast, our  
1056 results are in close agreement with the field-based measurements of Tang et al. (2020), which also  
1057 used a dynamic chamber flux method.  
1058

1059 These in-lab experiments show that simultaneous speciated  $N_r$  emissions [towards mass balance](#)  
 1060 [analysis](#) in a controlled environment can be conducted using a single chamber and potentially  
 1061 applied to an array of chambers, as others have done for a subset of gaseous  $N_r$  (Scharko et al.,  
 1062 2015; Tang et al., 2019). With [even-the](#) limited replicates we [did-explored](#) here, our results raise a  
 1063 question for researchers who have been using lab studies as a standard to predict and incorporate  
 1064 HONO soil emission values in particular; into regional and global models.

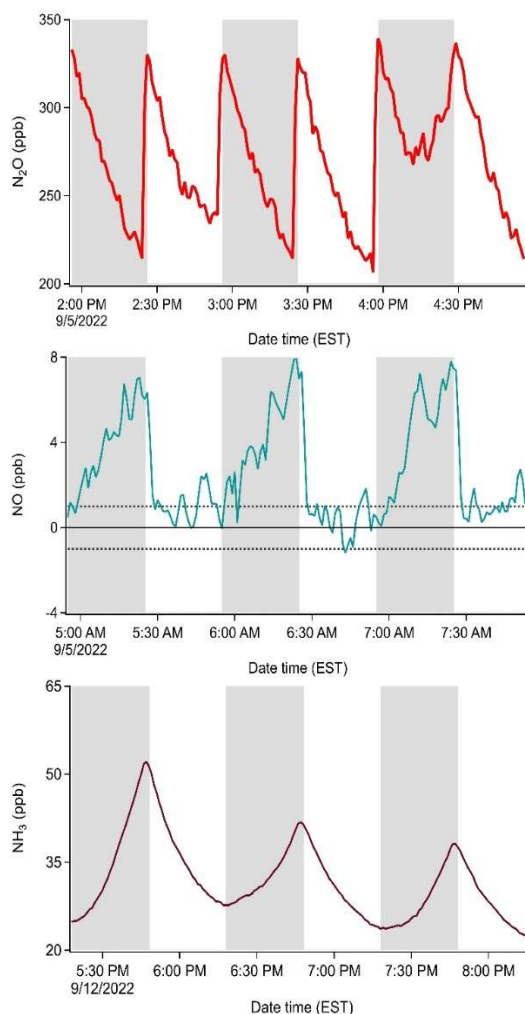


1065  
 1066 **Figure 4.** Relative flux contribution of soil treated with four different nitrogen-containing fertilizer  
 1067 salts. Each segment shows the proportion of integrated flux of NO (cyan), NO<sub>2</sub> (purple), HONO  
 1068 (blue), NH<sub>3</sub> (brown) and N<sub>2</sub>O (red), with discrete flux values presented in white text, and the total  
 1069 flux provided in black text above the column ( $\mu\text{g N m}^{-2}$ ). Some samples (denoted by \*) were only  
 1070 characterized for fluxes of NO, NO<sub>2</sub> and HONO using the modified NO<sub>x</sub> analyzer, as the duration  
 1071 of our access to the G2509 for NH<sub>3</sub> and N<sub>2</sub>O measurements was limited.

1072  
 1073 **3.3.3 Dual chamber field deployment for automated continuous dynamic fluxes**

1074 A pilot scale field campaign was carried out to demonstrate the application of our dual soil flux  
 1075 chambers in capturing  $N_r$  gas exchange processes; [taking into account the characterized processes](#)  
 1076 [presented above that are necessary for accurate determinations](#). A paired [measurement-MC](#) and  
 1077 [reference-chamber-RC](#) setup was deployed in the same field a year after the soil samples were

1078 collected for our lab experiments. During a period of stimulated  $N_r$  emission from an experimental  
1079 application of urea in situ, the mixing ratios of  $NO$ ,  $N_2O$ , and  $NH_3$  were impacted compared to the  
1080 unfertilized state. The changes within both the measurement-MC and reference-chambers-RC were  
1081 measured and used to calculate fluxes (Figure 5). The purpose of Figure 5 is methodological to  
1082 demonstrate how rate, dilution, and reaction terms combine in the observed rates of concentration  
1083 change during chamber cycles. The selected data for  $NH_3$  correspond to measurements taken after  
1084 the fertilization event, while the selected  $NO$  and  $N_2O$  data segments are examples for separate  
1085 observations times which best demonstrated the contributions of all terms prior to the fertilization  
1086 event. ~~so that~~ Taken together, these three separate examples for the mathematical terms in E6 and  
1087 E7 contributing to the net flux can be considered more easily – they are entirely ascribed to the  
1088 rate term otherwise. Each set of observations span three consecutive hours of dynamic changes in  
1089 gas concentrations within the chambers which allow fluxes to be calculated. manco measurement  
1090 rate of the Picarro, it is clear from the accumulation and depletion of target gases in Figure 5, that  
1091 a shorter observation period than 30 minutes could be used when high time resolution  
1092 instrumentation across all target species is available. The benefit of this would be to reduce both  
1093 the alteration of the composition of the chamber headspace and diverging physical conditions  
1094 between the chamber and ambient environment; ultimately obtaining better flux estimates.  
1095 However, for this pilot study, the 1-minute time resolution of the  $NO_x$  analyzer and the method for  
1096 determining HONO by difference with an annular denuder every 5 minutes required the 30-minute  
1097 interval. A shorter closure period could also have the drawback of worse flux detection limits when  
1098 fluxes are small, and more variability due to a less robust regression of the accumulation or  
1099 depletion rate. For example, this would increase the value of  $\lambda$  for  $NH_3$  (Figure S11; Section S7)  
1100 and its relative error (4% for a clean system, Section S7.3) as well as other surface-interacting  
1101 gases.



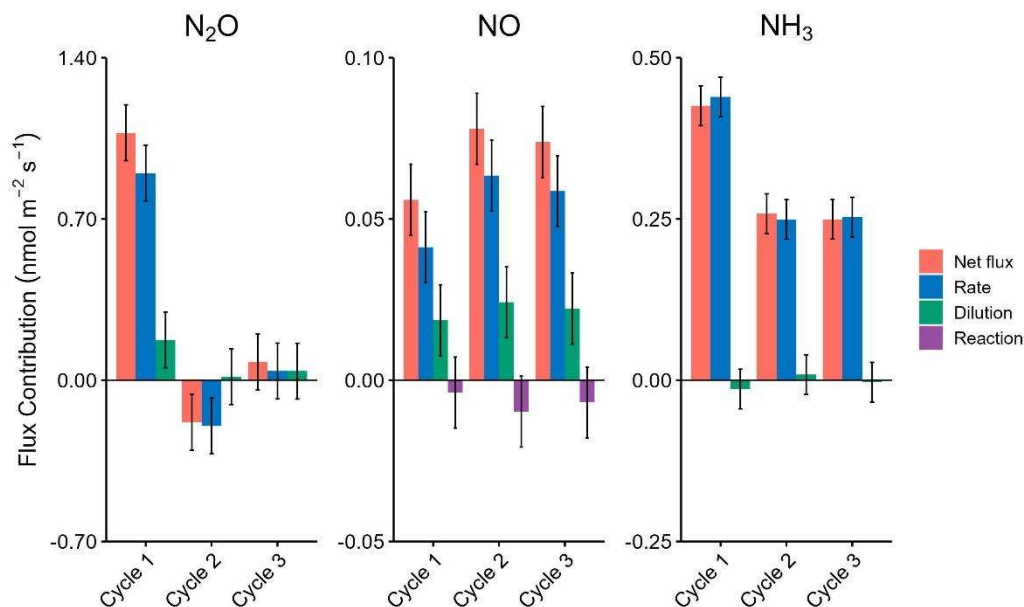
1102

1103 **Figure 5.** Mixing ratios of N<sub>2</sub>O (ppb), NO (ppb), and NH<sub>3</sub> (ppb) in the measurement (grey shading)  
 1104 and reference (unshaded) chambers from three consecutive cycles during the pilot field study.  
 1105 Cycles are shown for representative, non-simultaneous periods to illustrate term contributions.  
 1106 The NH<sub>3</sub> cycles correspond to measurements taken after the fertilization perturbation, while the  
 1107 NO and N<sub>2</sub>O cycles are from the same period before fertilization. The dashed lines for the NO  
 1108 measurements indicate ±1 ppb (3σ noise), where the positive boundary represents the detection  
 1109 limit of the NO<sub>x</sub> analyzer.

1110 The breakdown of the flux components for NH<sub>3</sub>, N<sub>2</sub>O, and NO can allow the contributions of the  
 1111 experimental setup (e.g. dilution) and environmental factors (e.g. reaction) to be considered  
 1112 independently (Figure 6). We consider these across a consecutive triplicate of flux determinations  
 1113 from the pilot field deployment, to provide meaningful examples. The uncertainty in each term is  
 1114 estimated from the variance between the triplicate of consecutive MC and reference-chamber RC

1115 ~~observations, through the~~ measured slopes (i.e. using E6 or E7) and assumes that ambient  
1116 atmospheric composition did not change substantially during this error derivation period.

1117 The rate term ( $dC/dt$ ) is the dominant contributor to the determined flux for all three gases (Figure  
1118 6), which are all positive and indicate emissions despite their clearly differing temporal trends.  
1119 Each contribution term is defined and discussed in further detail in Section S7 ~~of the~~  
1120 ~~Supplementary Information~~. Accurate quantification of trace gas fluxes using dynamic chamber  
1121 systems requires correction for attenuation caused by surface interactions. These effects can  
1122 significantly suppress the measured accumulation rate of reactive species within the closed  
1123 ~~chamberMC, particularly for compounds with notable surface affinities such as NH<sub>3</sub>, NO<sub>2</sub>, and~~  
1124 ~~HONO~~ (Figure 5). To account for ~~such losses~~ ~~the~~, an ~~the~~ attenuation factor ( $\lambda$ ) was introduced in  
1125 this work (ES12), defined as the ratio of the theoretical concentration signal ~~for an inert gas (e.g.~~  
1126 ~~as in our fill/empty experiments)~~ to the measured signal over the chamber closure period. The  
1127 value of  $\lambda$  is gas-specific and time-dependent, reflecting wall affinity and kinetics, ~~which are~~  
1128 ~~therefore calculated over the first 30 min of chamber closure, as this was the duration of our field~~  
1129 ~~measurement closures~~. This correction reduces ~~flux~~ bias ~~that would otherwise result in the~~  
1130 ~~calculated gas exchange rates~~, independent of chamber-specific losses. The highest attenuation  
1131 was observed for NH<sub>3</sub>, for which a  $\lambda$  of 5.40 was determined ~~empirically~~ for the ~~corresponding 30-~~  
1132 ~~minute~~ closure period. ~~That is, substantial bias would result without this correction.~~ Comparative  
1133 values for NO<sub>2</sub> and HONO are also provided in ~~the Supporting Information~~ (Section S7.3).



1134

1135 **Figure 6.** Contributions of different terms (rate, dilution, and reaction) to the net flux of NO, NH<sub>3</sub>,  
 1136 and N<sub>2</sub>O across three consecutive cycles (nmol m<sup>-2</sup> s<sup>-1</sup>). The rate represents the change in  
 1137 concentration measured over time, the dilution represents the integrated loss to dilution, and the  
 1138 reaction represents the contribution of known reactions happening inside the chamber. Error bars  
 1139 are calculated using the mean value standard deviation about the mean of the corresponding terms  
 1140 from three reference chambers plus their standard deviation consecutive chamber cycles. For visual  
 1141 clarity, the attenuation correction ( $\lambda = 5.40$ ) for NH<sub>3</sub> is not applied in this figure, as it constitutes  
 1142 a linear scaling factor across all NH<sub>3</sub> bars (Section S7.3).

1143 For N<sub>2</sub>O, the first cycle exhibits the highest flux ( $1.07 \pm 0.12$  nmol m<sup>-2</sup> s<sup>-1</sup>), while the second cycle  
 1144 showed uptake by the soil ( $-0.18 \pm 0.12$  nmol m<sup>-2</sup> s<sup>-1</sup>), and the final cycle showed no net flux within  
 1145 error ( $0.08 \pm 0.12$  nmol m<sup>-2</sup> s<sup>-1</sup>). This occurs for a non-reactive greenhouse gas like N<sub>2</sub>O despite the  
 1146 concentration decreasing with time for both the measurement and reference periods, as the rate of  
 1147 decrease during the measurement cycle is slower than that from dilution alone (Figure 6). Across  
 1148 the three cycles, the rate of concentration change term contributed  $84 \pm 15\%$  to the measured flux  
 1149 in the first cycle,  $109 \pm 98\%$  in the second cycle, and  $50 \pm 171\%$  in the third cycle. The negative rate  
 1150 of concentration change in all of the observation periods arises due to the use of N<sub>2</sub>O-free purge  
 1151 gas delivered to the chamber headspace, which is required for the gas sample to be destructively  
 1152 quantified, while not introducing ambient air into the chamber (Figure 5). This simultaneously  
 1153 prevents sudden changes in local outdoor air composition from making small fluxes difficult to  
 1154 detect, as well as reduces the uncertainty in the fluxes assigned. Here, the dilution terms range

1155 from 8 to 50% of the net flux, as would be expected from the N<sub>2</sub>O exchange switching from  
1156 emission to deposition during the three-cycle example (i.e. 3 hours). In the final measurement  
1157 cycle, the rate of concentration change transitions from a loss to steady state during the observation  
1158 period, suggesting that an instant of ‘hot spot, hot moment’ emission of N<sub>2</sub>O was likely occurring.  
1159 As such, the transition leads to no net flux for the observation period, which is a limitation of our  
1160 dual-chamber approach compared to one with recirculating headspace, or that of an eddy  
1161 covariance approach that can capture ~~much~~ higher temporal resolution changes in concentration.  
1162 The N<sub>2</sub>O fluxes observed here are consistent with those reported for European arable soils using  
1163 both dynamic and static chambers, where event-driven N<sub>2</sub>O peaks after fertilization commonly  
1164 range between 0.2 and 2.4 nmol m<sup>-2</sup> s<sup>-1</sup>, and background periods can be as low as 0.08 nmol m<sup>-2</sup> s<sup>-1</sup>  
1165 <sup>1</sup> (Kong et al., 2025, Murphy et al., 2022, Maier et al., 2024, [Manco et al., 2025](#)). Field-scale eddy  
1166 covariance studies, such as Maier et al. (2022), have captured post-harvest pulses up to 1.6 nmol m<sup>-2</sup>  
1167 s<sup>-1</sup>, highlighting the capacity of EC to resolve large, rapid emission events that single-point  
1168 chamber systems may miss. However, dynamic chamber systems provide high-precision,  
1169 temporally resolved flux data under controlled conditions, enabling direct attribution of emissions  
1170 to specific soil management or environmental factors (Butterbach-Bahl et al., 2013; Kong et al.,  
1171 2025). This makes dynamic chambers, like those demonstrated here, especially valuable for  
1172 mechanistic and process-level studies.

1173 For NO, the only gas we consider in the examples here with a reaction term, the reaction  
1174 contribution is the smallest among the three flux components. Over all three cycles, the reaction  
1175 term opposed the net flux by less than 13%. The magnitude of the reaction term is always  
1176 negligible, remaining within the ±0.01 nmol m<sup>-2</sup> s<sup>-1</sup> variability caused by background correction  
1177 from the ~~reference chamber~~ RC. Instead, the time rate of change in concentration term drove 74–  
1178 81% of the total flux for all three cycles and dilution accounted for the remaining 30–33%. The  
1179 reaction term falling within the uncertainty range of no contribution in all three cycles (e.g. -  
1180 0.01±0.01 nmol m<sup>-2</sup> s<sup>-1</sup> in the second cycle, where it had the greatest potential), indicates that its  
1181 contribution is less certain than the other flux components as it is driven by the presence of O<sub>3</sub>,  
1182 which is rapidly lost upon chamber closure. Overall, the low contribution of the reaction term  
1183 suggests this process has a minor role in measured NO fluxes. In other, more polluted regions,  
1184 such as the North China Plain (NCP) and the Pearl River Delta (PRD), where summertime ambient  
1185 O<sub>3</sub> concentrations ~~frequently~~ range from 60 to 275 ppbv, ~~and the~~ exceedances above 200 ppbv

1186 ~~occur~~ during pollution episodes (Wang et al., 2017); ~~this term~~ could ~~become~~ make this term very  
1187 important ~~to include~~. As noted above, the PFA film on the chamber surface will change its  
1188 properties with respect to O<sub>3</sub> transmission over time, highlighting the utility of the reference  
1189 chamber-RC, which is designed to accumulate the same atmospheric compounds as the  
1190 measurement chamber-MC over time on all sampling surfaces.

1191 The resulting emission fluxes of NO observed in these three cycles were 0.056-0.078 nmol m<sup>-2</sup> s<sup>-</sup>  
1192 <sup>1</sup>, which ~~is~~ are well within the range reported for agricultural soils in North America and Europe,  
1193 such as Taylor et al. (1999) who observed -0.07-4.2 nmol m<sup>-2</sup> s<sup>-1</sup> in Canadian fertilized cropland,  
1194 and Pape et al. (2009) and Almand-Hunter et al. (2014) who reported values of 0.05-4.0 nmol m<sup>-</sup>  
1195 <sup>2</sup> s<sup>-1</sup>, using dynamic chambers in grass and cropland soils. Therefore, in the case of NO for this  
1196 example, the variability in the fluxes is largely driven by real fluctuations in the rate term. This  
1197 highlights the sensitivity of the total flux to changes in the rate of concentration change, and the  
1198 precision of the method, as the uncertainty in the final fluxes here is on the order of 14%. Compared  
1199 to the dynamic chambers reported by these prior studies, ~~our the reference chamber-RC~~ in our  
1200 dual-chamber system offers a clear advantage by directly correcting for baseline fluctuations and  
1201 environmental drift that can confound single-chamber approaches. This is especially important for  
1202 reactive gases such as NO, which are susceptible to rapid loss to O<sub>3</sub> and short-term background  
1203 variability (Taylor et al., 1999). The dual-chamber system provides more robust, artifact-free  
1204 quantification of soil-atmosphere exchange and is less susceptible to interference from transient  
1205 local emissions (e.g. from nearby traffic or agricultural equipment) than single-chamber systems.  
1206 In comparison to eddy covariance or flux-gradient techniques, which integrate over larger areas  
1207 but may underestimate true NO fluxes due to post-emission chemistry (Taylor et al., 1999; Plake  
1208 et al., 2015), our approach yields high-frequency, process-resolving data ideal for mechanistic and  
1209 plot-scale studies.

1210 For NH<sub>3</sub>, the rate of change term dominates the total flux, contributing 103 ± 10%, 97 ± 16%, and  
1211 101 ± 17% of the flux in cycles one, two, and three, respectively, while the dilution term  
1212 contributes inconsequentially at -3 ± 7%, 3 ± 12%, and -1 ± 12% in the same three cycles,  
1213 respectively. The resulting emission fluxes of NH<sub>3</sub> observed across these cycles ranged from  
1214 0.43 ± 0.03 nmol m<sup>-2</sup> s<sup>-1</sup> in cycle one down to 0.25 ± 0.03 nmol m<sup>-2</sup> s<sup>-1</sup> in cycle three. The relatively  
1215 small contribution of the dilution term is consistent with the expectation that the reference-RC and

1216 ~~measurement chambers~~ MC exhibit similar dilution effects. The relative uncertainty in the final  
1217  $\text{NH}_3$  fluxes is 7% for cycle one, and 12% for cycles two and three. By comparison, relative  
1218 uncertainties were 14% for  $\text{NO}$  and ranged from 15% to 171% for  $\text{N}_2\text{O}$ , reflecting the greater  
1219 variability and lower precision associated with the smaller flux magnitudes for those gases.

1220 Our observed  $\text{NH}_3$  fluxes ( $0.25\text{--}0.43 \text{ nmol m}^{-2} \text{ s}^{-1}$ ) are consistent with literature values for managed  
1221 grass and croplands. For instance, Milford et al. (2009) reported bi-directional background fluxes  
1222 from  $-3.8$  to  $2.5 \text{ nmol m}^{-2} \text{ s}^{-1}$  prior to cutting intensively managed grassland, with larger diurnal  
1223 emissions up to  $42 \text{ nmol m}^{-2} \text{ s}^{-1}$  after cutting and maxima up to  $224 \text{ nmol m}^{-2} \text{ s}^{-1}$  following fertilizer  
1224 application. Abdulwahab et al. (2024) observed highly variable fluxes in intensively grazed French  
1225 grassland, ranging from  $-6.6$  to  $188 \text{ nmol m}^{-2} \text{ s}^{-1}$ , with short-lived maxima above  $300 \text{ nmol m}^{-2} \text{ s}^{-1}$   
1226 after slurry application, though most measurements were at much lower magnitudes. Notably,  
1227 accurate quantification of  $\text{NH}_3$  fluxes in this system critically depends on the application of the  
1228 chamber-specific attenuation factor  $\lambda$  (Section S.7.3; [Figure S11](#)). ~~In our study,  $\lambda$  for  $\text{NH}_3$  was 5.40~~  
1229 ~~for the 30-minute closure interval, meaning that w~~Without this correction, true  $\text{NH}_3$  fluxes would  
1230 be underestimated by more than fivefold ~~in our study~~. ~~The necessity of such a large  $\lambda$  correction is~~  
1231 ~~due to the strong surface affinity of  $\text{NH}_3$  and emphasizes the importance of it for gas and chamber-~~  
1232 ~~specific corrections to obtain high-quality  $\text{N}_r$  fluxes.~~ While the ~~reference chamber~~ RC correction  
1233 plays a role in the flux calculations, its impact on  $\text{NH}_3$  is less pronounced than on  $\text{N}_2\text{O}$ , where the  
1234 correction significantly alters the net flux direction (Figure 5; first cycle). Chamber-based  
1235 approaches such as ours provide a key advantage for  $\text{NH}_3$  over micrometeorological methods like  
1236 eddy covariance, which are especially prone to high-frequency attenuation and chemical  
1237 interferences for reactive gases. As shown in Moravek et al. (2019), even state-of-the-art closed-  
1238 path EC systems may recover less than half (as little as 46%) of true  $\text{NH}_3$  fluxes due to instrument  
1239 limitations and turbulence losses. Challenges were also evident for relaxed eddy accumulation  
1240 (REA), where Xu et al. (2010) found that turbulence and surface effects complicated flux  
1241 interpretation in cropland. Related methodological considerations have also been noted in other  
1242 contexts. Schlossberg et al. (2017) highlighted how airflow and canopy structure can influence  
1243 chamber  $\text{NH}_3$  fluxes in turf systems, underscoring the need for chamber methods that minimize  
1244 such artifacts. In contrast, our dynamic chamber system, with a chamber-specific  $\lambda$  correction and  
1245 ~~reference chamber~~ RC, enables robust, bias-corrected quantification of both low and episodic  $\text{NH}_3$   
1246 fluxes, as well as clear partitioning of emission and dilution terms, even under highly variable field

1247 conditions. Overall, the combination of dual-chamber design with explicit  $\lambda$  correction in our  
1248 method provides more accurate and robust quantification of soil  $\text{NH}_3$  emissions than single-  
1249 chamber, eddy covariance, or relaxed eddy accumulation techniques, particularly for short-term or  
1250 low-flux events.

#### 1252 4. Conclusions

1253 In this work, we presented a dynamic chamber system for [reactive nitrogen- \$\text{N}\_r\$](#)  flux measurements,  
1254 developed for the first time through the modification of commercially available chambers by  
1255 implementing two key changes: the use of PTFE fittings instead of original brass fittings and the  
1256 installation of an inert PFA film, which retains their actinic transparency. These modifications  
1257 provide a targeted methodology for other researchers to convert commercially available chambers  
1258 into those capable of measuring  $\text{N}_r$ . The performance of these modifications was quantified  
1259 through the rise and fall time constants of target gas concentrations, as well as a reduction in  
1260 reactive losses. [The time constants for the transfer of GHGs were not different from those of a](#)  
1261 [theoretically inert gas and showed no change from the modifications. Improved transmission for](#)  
1262 [the reactive and surface-active  \$\text{N}\_r\$  species  \$\text{NO}\_2\$ , HONO, and  \$\text{NH}\_3\$  targeted here ranged from 0.8 to](#)  
1263 [2.6 minutes. \[Further improvement for  \\$\text{NH}\\_3\\$  might be obtained by integration of heated gas transfer\]\(#\)](#)  
1264 [lines between the chambers and gas analyzers.](#) Similarly, a commercial chamber multiplexing unit  
1265 [with stainless steel valves and fittings replaced with PTFE and PFA, respectively, resulted in a 9–](#)  
1266 [27% reduction in surface losses of  \$\text{N}\_r\$  compounds.](#)

1267 Only  $\text{NO}_2$  showed reactive loss in the system, and the loss fraction in the chambers in their  
1268 unmodified configurations was up to  $\sim 36\%$ . Losses were reduced to the gas analyzer detection  
1269 limits ( $<10\%$ ) with the same fluoropolymer modifications when atmospherically relevant  $\text{NO}_2$  and  
1270 RH mixtures were introduced. Lastly,  $\text{O}_3$  loss was pervasive within chambers in both their  
1271 modified and unmodified configurations at 35% and 45% respectively, demonstrating the  
1272 necessity of the [reference-chamber-RC](#). It allows characterization of deposited surface film  
1273 reactivity for  $\text{O}_3$ , especially when obtaining quality  $\text{O}_3$  measurements is needed to account for the  
1274 conversion of  $\text{NO}$  to  $\text{NO}_2$  during a sampling period. Taken together, the modified commercial  
1275 system is capable of measuring dynamic soil fluxes of GHGs and  $\text{N}_r$  when [a measurement-MC](#) and  
1276 [reference-chamber-RC](#) are deployed simultaneously.

1277 Proof of concept flux measurements using this  $N_r$  dynamic chamber system were conducted using  
1278 real agricultural soil samples in the lab, with a single chamber, and during a pilot study in the field  
1279 with a measurement-reference chamber pair. The lab soil emissions were found to be consistent  
1280 with prior field reports in the literature for NO and HONO, with unexpected emissions of  $NO_2$  also  
1281 observed. Substantial variability in  $NO_2$  and HONO emissions between replicates demonstrates  
1282 the potential heterogeneity of soil emissions that may result from samples kept intact and subject  
1283 to minimal preparation conditions. Upon the addition of typical fertilizers, like urea, substantial  
1284  $NH_3$  and  $N_2O$  emissions fluxes matched expectations, with increasing quantities of NO,  $NO_2$ , and  
1285 HONO as the soil water content decreased.

1286 Last, fully automated operation of the chambers was carried out in a field pilot study with the  
1287 delivery of external fertilizer conducted at the midpoint of the [study two-week period](#) to stimulate  
1288  $N_r$  emissions, mainly in the form of  $NH_3$ . Continuous flux observations were made by switching  
1289 sample flows between ~~a paired reference and measurement chamber~~ [the RC and MC](#) with a  
1290 custom-built valve system to obtain continuous  $N_r$  fluxes ~~over two weeks~~. While the details of the  
1291 entire campaign will be presented in a future manuscript, it was shown here how the [developed](#)  
1292 ~~methodology developed measurement technique~~ yields reliable flux determinations by accounting  
1293 for known reactions and our lab-derived surface effects. The mathematical foundation of each term  
1294 and its error in a mass balance equation, which are central to reducing flux bias, are fully described.  
1295 The observed fluxes from our dynamic chambers are simpler to quantify with standard gas  
1296 analyzers and can be compared readily with prior reports in the literature from other common flux  
1297 techniques. [The duration of the chamber closure for the modified  \$NO\_x\$  analyzer used in this work  
1298 made a 30-minute observation necessary to obtain sufficient measurements for a reliable flux  
1299 determination. Use of higher time resolution gas analyzers can readily reduce the observation  
1300 period and continue to produce reliable fluxes, depending on the flux detection limit desired, and  
1301 tolerance for inflated surface effects.](#) One key limitation of the dynamic chamber approach is the  
1302 limited footprint covered in a system that is notoriously heterogeneous, where [EC eddy covariance](#)  
1303 and [REA relaxed eddy accumulation](#) flux approaches are less prone to this effect. Using a  
1304 multiplexer and a larger array of dynamic chambers, it is possible to reduce the susceptibility of  
1305 the dynamic chamber approach to this issue. [One further limitation is that the chamber and gas  
1306 transfer line surfaces may accumulate material and change the mass transfer of gases, particularly  
1307  \$NH\_3\$ , over time. This effect can be monitored using the reference chamber if sufficient quantities](#)

1308 are present in the ambient air, via the dilution decay rate and characterized in situ by a standard  
1309 addition of the necessary gases to the headspace.

1310 Through our modifications and validation, this work provides insight into how commercial  
1311 dynamic chamber options, like those offered by Eosense, can be modified easily for scientists from  
1312 various disciplines interested in studying  $N_r$  exchange at atmospheric interfaces. This is  
1313 particularly important for research groups which currently do not have the expertise and resources  
1314 to develop their own  $N_r$  flux measurement systems. The modified system utilizes destructive  
1315 sampling techniques, as opposed to the re-circulation of chamber air, enabling integration of the  
1316 dynamic chamber approach with various standard gas analysis instruments (e.g.,  $NO_x$   
1317 chemiluminescence) to study their exchange. The large footprint allows gas concentrations to  
1318 change in easily determined quantities even for very small fluxes. We show here, for the first time,  
1319 that such a system can provide simultaneous measurement of  $NO$ ,  $NO_2$ ,  $HONO$ ,  $NH_3$ ,  $CO_2$ ,  $H_2O$ ,  
1320  $CH_4$ , and  $N_2O$  fluxes. In addition, we show fully automated operation of the chambers, switching  
1321 of sample flows, and data collection workflows for continuous and unattended measurements of  
1322 fluxes at field sites. Ideally, modern instrumentation like the ToF-CIMS would be coupled with  
1323 this system to shorten the chamber closure duration, and better distinguish  $HONO$  from other  $NO_y$   
1324 fluxes, instead of a modified  $NO_x$  analyzer. Given the pressing need to understand global  
1325 perturbations to the biogeochemical cycle of N, reduce nitrogen use in agriculture, and gauge the  
1326 impacts of N status on biodiversity or ecosystem function, wider accessibility of  $N_r$  flux techniques  
1327 for the global research community is needed to increase the pace of research outcomes and improve  
1328 the capacity for interdisciplinary work between atmospheric and earth system researchers.

1329

1330 **Data Availability:** Data is available upon request from the Corresponding Author.

1331 **Competing Interests:** TCV received supporting in-kind funds for this work from Eosense, Inc.  
1332 and Picarro as it is mandatory in the NSERC Alliance Missions programme funding structure  
1333 which facilitates research partnerships between the academy and industry. NN, CC, and SE are  
1334 employed by Eosense, Inc.

1335 **Author contributions:** TCV designed and oversaw the experiments, acquired funding, wrote parts  
1336 of the manuscript, and guided writing and revision of all sections of the manuscript. MS wrote the

1337 manuscript and performed all the lab experiments. KZA carried out the multiplexer experiments,  
1338 conducted data analysis for the fill-empty experiment, provided guidance and feedback on the  
1339 derivation of the mass balance flux model, and assisted in the revision of the manuscript. DF  
1340 conducted some of the field measurements, worked up the pilot field study results and derived the  
1341 dual chamber mass balance flux calculations, and made contributions to the writing and revision  
1342 of the manuscript. LRC helped design the experiments, modify the chambers, write the LabVIEW  
1343 code, conduct the pilot field measurements, and contributed to manuscript preparation and  
1344 revision. AM contributed to the initial chamber lab setup and revision of the manuscript. FS  
1345 performed some fill-empty and NO<sub>2</sub> loss experiments, designed the custom valve switching  
1346 system, modified LabVIEW code for the pilot field measurements, and contributed to initial drafts  
1347 of the manuscript. YEI and TH assisted with the GHG fill-empty experiments, YEI supported the  
1348 pilot field study, and both contributed to manuscript revision. NN, CC, and SE provided technical  
1349 support with setting up and modifying the chambers and multiplexer.

1350 **Acknowledgements:** We gratefully acknowledge the support of Picarro for facilitating use of the  
1351 G2509 analyzer during the chamber characterizations, lab experiments, and pilot study  
1352 components of this work. The team at Shawnasey Farms Ltd. provided access to storage and  
1353 machinery to install the experimental agricultural site for the pilot project, helped with daily  
1354 campaign logistics, and the collection of soil samples.

1355 **Financial Support:** All components of this work were supported by funding to TCV by the  
1356 Natural Sciences and Engineering Research Council (NSERC) Alliance Missions program  
1357 (ALLRP 570577-2021), with additional support from the NSERC Discovery Grants and Early  
1358 Career Launch programmes (RGPIN-2020-06166 and DGEGR-2020-00186). MS and YEI were  
1359 supported by Ontario Graduate Scholarships, and MS was further supported by a Charles Hantho  
1360 Award in Atmospheric Chemistry and the Enbridge Graduate Student Award.

1361

1362 **5. References**

- 1363 Almand-Hunter, B. B., Walker, J. T., Masson, N. P., Hafford, L., & Hannigan, M. P. (2015).  
1364 Development and validation of inexpensive, automated, dynamic flux chambers.  
1365 *Atmospheric Measurement Techniques*, 8(1), 267–280. [https://doi.org/10.5194/amt-8-267-](https://doi.org/10.5194/amt-8-267-2015)  
1366 [2015](https://doi.org/10.5194/amt-8-267-2015)
- 1367 Anas, M., Liao, F., Verma, K. K., Sarwar, M. A., Mahmood, A., Chen, Z. L., Li, Q., Zeng, X. P.,  
1368 Liu, Y., & Li, Y. R. (2020). Fate of nitrogen in agriculture and environment: agronomic,  
1369 eco-physiological and molecular approaches to improve nitrogen use efficiency. In  
1370 *Biological Research* (Vol. 53, Issue 1). BioMed Central Ltd.  
1371 <https://doi.org/10.1186/s40659-020-00312-4>
- 1372 Aneja, V. P., Blunden, J., Claiborn, C. S., & Rogers, H. H. (2006). Dynamic Chamber System to  
1373 Measure Gaseous Compounds Emissions and Atmospheric-Biospheric Interactions.  
1374 *Environmental Simulation Chambers: Application to Atmospheric Chemical Processes*, 97–  
1375 109. [https://doi.org/10.1007/1-4020-4232-9\\_7](https://doi.org/10.1007/1-4020-4232-9_7)
- 1376 Anthony, T. L., & Silver, W. L. (2024). Hot spots and hot moments of greenhouse gas emissions  
1377 in agricultural peatlands. *Biogeochemistry*, 167(4), 461–477.  
1378 <https://doi.org/10.1007/s10533-023-01095-y>
- 1379 Bao, F., Cheng, Y., Kuhn, U., Li, G., Wang, W., Kratz, A. M., Weber, J., Weber, B., Pöschl, U.,  
1380 & Su, H. (2022). Key Role of Equilibrium HONO Concentration over Soil in Quantifying  
1381 Soil–Atmosphere HONO Fluxes. *Environmental Science & Technology*.  
1382 <https://doi.org/10.1021/acs.est.1c06716>
- 1383 Barney, W. S., & Finlayson-Pitts, B. J. (2000). Enhancement of N<sub>2</sub>O<sub>4</sub> on Porous Glass at Room  
1384 Temperature: A Key Intermediate in the Heterogeneous Hydrolysis of NO<sub>2</sub>? *Journal of*  
1385 *Physical Chemistry A*, 104(2), 171–175. <https://doi.org/10.1021/jp993169b>
- 1386 Becciolini, V., Leso, L., Fuertes Gimeno, E., Rossi, G., Barbari, M., Dalla Marta, A., Orlandini,  
1387 S., & Verdi, L. (2024). Nitrogen loss abatement from dairy cow excreta through urine and  
1388 faeces separation: The effect of temperature and exposure period on NH<sub>3</sub> fluxes.  
1389 *Agricultural Systems*, 216. <https://doi.org/10.1016/j.agsy.2024.103898>
- 1390 Behera, S. N., Sharma, M., Aneja, V. P., & Balasubramanian, R. (2013). Ammonia in the  
1391 atmosphere: A review on emission sources, atmospheric chemistry and deposition on  
1392 terrestrial bodies. In *Environmental Science and Pollution Research* (Vol. 20, Issue 11, pp.  
1393 8092–8131). <https://doi.org/10.1007/s11356-013-2051-9>
- 1394 Behrendt, T., Veres, P. R., Ashuri, F., Song, G., Flanz, M., Mamtimin, B., Bruse, M., Williams,  
1395 J., & Meixner, F. X. (2014). Characterisation of NO production and consumption: New  
1396 insights by an improved laboratory dynamic chamber technique. *Biogeosciences*, 11(19),  
1397 5463–5492. <https://doi.org/10.5194/bg-11-5463-2014>
- 1398 Benedict, K. B., Prenni, A. J., Carrico, C. M., Sullivan, A. P., Schichtel, B. A., & Collett, J. L.  
1399 (2017). Enhanced concentrations of reactive nitrogen species in wildfire smoke.  
1400 *Atmospheric Environment*, 148, 8–15. <https://doi.org/10.1016/j.atmosenv.2016.10.030>
- 1401 Burkholder, J. B., Cox, R. A., & Ravishankara, A. R. (2015). Atmospheric Degradation of Ozone  
1402 Depleting Substances, Their Substitutes, and Related Species. *Chemical Reviews*, 115(10),  
1403 3704–3759. <https://doi.org/10.1021/cr5006759>

1404 Butterbach-Bahl, K., & Dannenmann, M. (2011). Denitrification and associated soil N<sub>2</sub>O  
1405 emissions due to agricultural activities in a changing climate. *Current Opinion in*  
1406 *Environmental Sustainability*, 3(5), 389–395. <https://doi.org/10.1016/j.cosust.2011.08.004>

1407 Chiaravalloti, I., Theunissen, N., Zhang, S., Wang, J., Sun, F., Ahmed, A. A., Pihlap, E.,  
1408 Reinhard, C. T., & Planavsky, N. J. (2023). Mitigation of soil nitrous oxide emissions  
1409 during maize production with basalt amendments. *Frontiers in Climate*, 5.  
1410 <https://doi.org/10.3389/fclim.2023.1203043>

1411 Crilley, L. R., Kramer, L. J., Ouyang, B., Duan, J., Zhang, W., Tong, S., Ge, M., Tang, K., Qin,  
1412 M., Xe, P., Shaw, M. D., Lewis, A. C., Mehra, A., Bannan, T. J., Worrall, S. D., Priestley,  
1413 M., Bacak, A., Coe, H., Allan, J., ... Bloss, W. J. (2019). *Intercomparison of nitrous acid*  
1414 *(HONO) measurement techniques in a megacity (Beijing)*. [https://doi.org/10.5194/amt-](https://doi.org/10.5194/amt-2019-139)  
1415 [2019-139](https://doi.org/10.5194/amt-2019-139)

1416 Crilley, L. R., Lao, M., Salehpoor, L., & VandenBoer, T. C. (2023). Emerging investigator  
1417 series: an instrument to measure and speciate the total reactive nitrogen budget indoors:  
1418 description and field measurements. *Environmental Science: Processes and Impacts*, 25(3),  
1419 389–404. <https://doi.org/10.1039/d2em00446a>

1420 Degaspari, I. A. M., Soares, J. R., Montezano, Z. F., Del Grosso, S. J., Vitti, A. C., Rossetto, R.,  
1421 & Cantarella, H. (2020). Nitrogen sources and application rates affect emissions of N<sub>2</sub>O  
1422 and NH<sub>3</sub> in sugarcane. *Nutrient Cycling in Agroecosystems*, 116(3), 329–344.  
1423 <https://doi.org/10.1007/s10705-019-10045-w>

1424 Delaria, E. R., & Cohen, R. C. (2023). Measurements of Atmosphere-Biosphere Exchange of  
1425 Oxidized Nitrogen and Implications for the Chemistry of Atmospheric NO<sub>x</sub>. *Accounts of*  
1426 *Chemical Research*, 56(13), 1720–1730. <https://doi.org/10.1021/acs.accounts.3c00090>

1427 Dennis, R. L., Mathur, R., Pleim, J. E., & Walker, J. T. (2010). Fate of ammonia emissions at the  
1428 local to regional scale as simulated by the Community Multiscale Air Quality model.  
1429 *Atmospheric Pollution Research*, 1(4), 207–214. <https://doi.org/10.5094/APR.2010.027>

1430 Ebnesajjad, S. (2017). Introduction to Fluoropolymers. *Applied Plastics Engineering Handbook:*  
1431 *Processing, Materials, and Applications: Second Edition*, 55–71.  
1432 <https://doi.org/10.1016/B978-0-323-39040-8.00003-1>

1433 Edwards, T. M., Puglis, H. J., Kent, D. B., Durán, J. L., Bradshaw, L. M., & Farag, A. M.  
1434 (2024). Ammonia and aquatic ecosystems – A review of global sources, biogeochemical  
1435 cycling, and effects on fish. In *Science of the Total Environment* (Vol. 907). Elsevier B.V.  
1436 <https://doi.org/10.1016/j.scitotenv.2023.167911>

1437 Ellis, R. A., Murphy, J. G., Pattey, E., Van Haarlem, R., O'Brien, J. M., & Herndon, S. C.  
1438 (2010). Characterizing a Quantum Cascade Tunable Infrared Laser Differential Absorption  
1439 Spectrometer (QC-TILDAS) for measurements of atmospheric ammonia. *Atmospheric*  
1440 *Measurement Techniques*, 3(2), 397–406. <https://doi.org/10.5194/amt-3-397-2010>

1441 Finlayson-Pitts, B. J., Wingen, L. M., Sumner, A. L., Syomin, D., & Ramazan, K. A. (2003). The  
1442 heterogeneous hydrolysis of NO<sub>2</sub> in laboratory systems and in outdoor and indoor  
1443 atmospheres: An integrated mechanism. *Physical Chemistry Chemical Physics*, 5(2), 223–  
1444 242. <https://doi.org/10.1039/b208564j>

1445 Fowler, D., Coyle, M., Skiba, U., Sutton, M. A., Cape, J. N., Reis, S., Sheppard, L. J., Jenkins,  
1446 A., Grizzetti, B., Galloway, J. N., Vitousek, P., Leach, A., Bouwman, A. F., Butterbach-  
1447 Bahl, K., Dentener, F., Stevenson, D., Amann, M., & Voss, M. (2013). The global nitrogen  
1448 cycle in the twenty-first century. *Philosophical Transactions of the Royal Society B:  
1449 Biological Sciences*, 368(1621). <https://doi.org/10.1098/rstb.2013.0164>

1450 Geddes, J. A., & Murphy, J. G. (2014). Observations of reactive nitrogen oxide fluxes by eddy  
1451 covariance above two midlatitude North American mixed hardwood forests. *Atmospheric  
1452 Chemistry and Physics*, 14(6), 2939–2957. <https://doi.org/10.5194/acp-14-2939-2014>

1453 George, C., Ammann, M., D’Anna, B., Donaldson, D. J., & Nizkorodov, S. A. (2015).  
1454 Heterogeneous Photochemistry in the Atmosphere. *Chemical Reviews*, 115(10), 4218–4258.  
1455 <https://doi.org/10.1021/cr500648z>

1456 Gong, C., Wang, Y., Tian, H., Kou-Giesbrecht, S., Vuichard, N., & Zaehle, S. (2025).  
1457 Uncertainties in fertilizer-induced emissions of soil nitrogen oxide and the associated  
1458 impacts on ground-level ozone and methane. *EGUsphere*.  
1459 <https://doi.org/10.5194/egusphere-2025-1416>

1460 González Ortiz, Alberto., Guerreiro, Cristina., & Soares, Joana. (2020). *Air quality in Europe:  
1461 2020 report*. European Environment Agency. <https://doi.org/10.2800/786656>

1462 Govoni Brondi, M., Bortoletto-Santos, R., Farias, J. G., Farinas, C. S., Ammar, M., Ribeiro, C.,  
1463 Williams, C., & Baltrusaitis, J. (2024). Mechanochemically Synthesized Nitrogen-Efficient  
1464 Mg- and Zn-Ammonium Carbonate Fertilizers. *ACS Sustainable Chemistry and  
1465 Engineering*, 12(16), 6182–6193. <https://doi.org/10.1021/acssuschemeng.3c07785>

1466 Ha, P. T. M., Kanaya, Y., Taketani, F., Andres Hernandez, M. D., Schreiner, B., Pfeilsticker, K.,  
1467 & Sudo, K. (2023). Implementation of HONO into the chemistry-climate model CHASER  
1468 (V4.0): Roles in tropospheric chemistry. *Geoscientific Model Development*, 16(3), 927–960.  
1469 <https://doi.org/10.5194/gmd-16-927-2023>

1470 He, Y., Zhou, X., Hou, J., Gao, H., & Bertman, S. B. (2006). Importance of dew in controlling  
1471 the air-surface exchange of HONO in rural forested environments. *Geophysical Research  
1472 Letters*, 33(2). <https://doi.org/10.1029/2005GL024348>

1473 Huang, G., Zhou, X., Deng, G., Qiao, H., & Civerolo, K. (2002). Measurements of atmospheric  
1474 nitrous acid and nitric acid. *Atmospheric Environment*, 36(13), 2225–2235.  
1475 [https://doi.org/10.1016/S1352-2310\(02\)00170-X](https://doi.org/10.1016/S1352-2310(02)00170-X)

1476 Huber, D. E., Kort, E. A., & Steiner, A. L. (2024). Soil Moisture, Soil NO<sub>x</sub> and Regional Air  
1477 Quality in the Agricultural Central United States. *Journal of Geophysical Research:  
1478 Atmospheres*, 129(12). <https://doi.org/10.1029/2024JD041015>

1479 IPCC. (2023). *Climate Change 2021- The Physical Science Basis*. Cambridge University Press.  
1480 <https://doi.org/https://doi.org/10.1017/9781009157896>

1481 Jang, J.-H., Hong, J., Kim, J. B., Park, S., Hwang, K., Kim, J., Kim, J. Y., Bae, G.-N., Kim, S., &  
1482 Kim, K. H. (2025). Influence of atmospheric ammonia on secondary inorganic aerosol  
1483 formation in PM<sub>2.5</sub> during spring 2024 in the Hongseong area, Republic of Korea.  
1484 *Atmospheric Environment*, 121363. <https://doi.org/10.1016/j.atmosenv.2025.121363>

1485 Kamboures, M. A., Raff, J. D., Miller, Y., Phillips, L. F., Finlayson-Pitts, B. J., & Gerber, R. B.  
1486 (2008). Complexes of HNO<sub>3</sub> and NO<sub>3</sub>- with NO<sub>2</sub> and N<sub>2</sub>O<sub>4</sub>, and their potential role in

1487 atmospheric HONO formation. *Physical Chemistry Chemical Physics*, 10(39), 6019–6032.  
1488 <https://doi.org/10.1039/b805330h>

1489 Kamp, J. N., Häni, C., Nyord, T., Feilberg, A., & Sørensen, L. L. (2020). The aerodynamic  
1490 gradient method: Implications of non-simultaneous measurements at alternating heights.  
1491 *Atmosphere*, 11(10). <https://doi.org/10.3390/atmos11101067>

1492 Kleffmann, J., Gavriloaiei, T., Hofzumahaus, A., Holland, F., Koppmann, R., Rupp, L.,  
1493 Schlosser, E., Siese, M., & Wahner, A. (2005). Daytime formation of nitrous acid: A major  
1494 source of OH radicals in a forest. *Geophysical Research Letters*, 32(5), 1–4.  
1495 <https://doi.org/10.1029/2005GL022524>

1496 Kolari, P., Bäck, J., Taipale, R., Ruuskanen, T. M., Kajos, M. K., Rinne, J., Kulmala, M., &  
1497 Hari, P. (2012). Evaluation of accuracy in measurements of VOC emissions with dynamic  
1498 chamber system. *Atmospheric Environment*, 62, 344–351.  
1499 <https://doi.org/10.1016/j.atmosenv.2012.08.054>

1500 Kool, D. M., Wrage, N., Zechmeister-Boltenstern, S., Pfeffer, M., Brus, D., Oenema, O., & Van  
1501 Groenigen, J. W. (2010). Nitrifier denitrification can be a source of N<sub>2</sub>O from soil: A  
1502 revised approach to the dual-isotope labelling method. *European Journal of Soil Science*,  
1503 61(5), 759–772. <https://doi.org/10.1111/j.1365-2389.2010.01270.x>

1504 Lao, M., Crilley, L. R., Salehpoor, L., Furlani, T. C., Bourgeois, I., Andrew Neuman, J., Rollins,  
1505 A. W., Veres, P. R., Washenfelder, R. A., Womack, C. C., Young, C. J., & VandenBoer, T.  
1506 C. (2020). A portable, robust, stable, and tunable calibration source for gas-phase nitrous  
1507 acid (HONO). *Atmospheric Measurement Techniques*, 13(11), 5873–5890.  
1508 <https://doi.org/10.5194/amt-13-5873-2020>

1509 Laufs, S., Cazaunau, M., Stella, P., Kurtenbach, R., Cellier, P., Mellouki, A., Loubet, B., &  
1510 Kleffmann, J. (2017). Diurnal fluxes of HONO above a crop rotation. *Atmospheric  
1511 Chemistry and Physics*, 17(11), 6907–6923. <https://doi.org/10.5194/acp-17-6907-2017>

1512 Lee, B. H., Lopez-Hilfiker, F. D., Mohr, C., Kurtén, T., Worsnop, D. R., & Thornton, J. A.  
1513 (2014). An iodide-adduct high-resolution time-of-flight chemical-ionization mass  
1514 spectrometer: Application to atmospheric inorganic and organic compounds. *Environmental  
1515 Science and Technology*, 48(11), 6309–6317. <https://doi.org/10.1021/es500362a>

1516 Lehnert, N., Musselman, B. W., Seefeldt, L. C., & Gutenberg-university, J. (2021). Grand  
1517 Challenges in the nitrogen cycle. *Chemical Society Reviews*, 50, 3640–3646.  
1518 <https://doi.org/10.1039/d0cs00923g>

1519 Li, L., Fan, W., Kang, X., Wang, Y., Cui, X., Xu, C., Griffin, K. L., & Hao, Y. (2016).  
1520 Responses of greenhouse gas fluxes to climate extremes in a semiarid grassland.  
1521 *Atmospheric Environment*, 142, 32–42. <https://doi.org/10.1016/j.atmosenv.2016.07.039>

1522 Lipiec, J., Walczak, R., Witkowska-Walczak, B., Nosalewicz, A., Słowińska-Jurkiewicz, A., &  
1523 Sławiński, C. (2007). The effect of aggregate size on water retention and pore structure of  
1524 two silt loam soils of different genesis. *Soil and Tillage Research*, 97(2), 239–246.  
1525 <https://doi.org/10.1016/j.still.2007.10.001>

1526 Liu, L., Zhang, X., Xu, W., Liu, X., Li, Y., Wei, J., Wang, Z., & Lu, X. (2020). Ammonia  
1527 volatilization as the major nitrogen loss pathway in dryland agro-ecosystems.  
1528 *Environmental Pollution*, 265. <https://doi.org/10.1016/j.envpol.2020.114862>

- 1529 Liu, Zheng, X., Li, Y., Yu, J., Ding, H., Sveen, T. R., & Zhang, Y. (2022). Soil moisture  
 1530 determines nitrous oxide emission and uptake. *Science of the Total Environment*, 822.  
 1531 <https://doi.org/10.1016/j.scitotenv.2022.153566>
- 1532 Ludwig, J., Meixner, F. X., Vogel, B., & Forstner, J. (2001). Soil-air exchange of nitric oxide:  
 1533 An overview of processes, environmental factors, and modeling studies. *Biogeochemistry*,  
 1534 52(3), 225–257. <https://doi.org/10.1023/A:1006424330555>
- 1535 Luo, X., Zhang, M., Ni, Y., & Shen, G. (2025). Mitigation strategies for NH<sub>3</sub> and N<sub>2</sub>O emissions  
 1536 in greenhouse agriculture: Insights into fertilizer management and nitrogen emission  
 1537 mechanisms. *Environmental Technology and Innovation*, 37.  
 1538 <https://doi.org/10.1016/j.eti.2024.103995>
- 1539 Maggiotto, S. R., Webb, J. A., & Thurtell, G. W. (2000). Nitrous and Nitrogen Oxide Emissions  
 1540 from Turfgrass Receiving Different Forms of Nitrogen Fertilizer. *American Society of*  
 1541 *Agronomy, Crop Science Society of America, and Soil Science Society of America*, 29, 621–  
 1542 630.
- 1543 Manco, A., Giaccone, M., Vitale, L., Maglione, G., Riccardi, M., Matteo, B. Di, Esposito, A.,  
 1544 Magliulo, V., & Tedeschi, A. (2025). Comparative Effects of Nitrogen Fertigation and  
 1545 Granular Fertilizer Application on Pepper Yield and Soil GHGs Emissions. *Horticulturae*,  
 1546 11(6). <https://doi.org/10.3390/horticulturae11060708>
- 1547 Mangalassery, S., Sjögersten, S., Sparkes, D. L., Sturrock, C. J., & Mooney, S. J. (2013). The  
 1548 effect of soil aggregate size on pore structure and its consequence on emission of  
 1549 greenhouse gases. *Soil and Tillage Research*, 132, 39–46.  
 1550 <https://doi.org/10.1016/j.still.2013.05.003>
- 1551 Meusel, H., Tamm, A., Kuhn, U., Wu, D., Lena Leifke, A., Fiedler, S., Ruckteschler, N.,  
 1552 Yordanova, P., Lang-Yona, N., Pöhlker, M., Lelieveld, J., Hoffmann, T., Pöschl, U., Su, H.,  
 1553 Weber, B., & Cheng, Y. (2018). Emission of nitrous acid from soil and biological soil crusts  
 1554 represents an important source of HONO in the remote atmosphere in Cyprus. *Atmospheric*  
 1555 *Chemistry and Physics*, 18(2), 799–813. <https://doi.org/10.5194/acp-18-799-2018>
- 1556 Min, K. E., Pusede, S. E., Browne, E. C., LaFranchi, B. W., & Cohen, R. C. (2014). Eddy  
 1557 covariance fluxes and vertical concentration gradient measurements of NO and NO<sub>2</sub> over a  
 1558 ponderosa pine ecosystem: Observational evidence for within-canopy chemical removal of  
 1559 NO<sub>x</sub>. *Atmospheric Chemistry and Physics*, 14(11), 5495–5512. <https://doi.org/10.5194/acp-14-5495-2014>
- 1561 Mochizuki, T., Amagai, T., & Tani, A. (2018). Effects of soil water content and elevated CO<sub>2</sub>  
 1562 concentration on the monoterpene emission rate of *Cryptomeria japonica*. *Science of the*  
 1563 *Total Environment*, 634, 900–908. <https://doi.org/10.1016/j.scitotenv.2018.04.025>
- 1564 Moravek, A., Foken, T., & Trebs, I. (2014). Application of a GC-ECD for measurements of  
 1565 biosphere-atmosphere exchange fluxes of peroxyacetyl nitrate using the relaxed eddy  
 1566 accumulation and gradient method. *Atmospheric Measurement Techniques*, 7(7), 2097–  
 1567 2119. <https://doi.org/10.5194/amt-7-2097-2014>
- 1568 Moravek, A., Singh, S., Pattey, E., Pelletier, L., & Murphy, J. G. (2019). Measurements and  
 1569 quality control of ammonia eddy covariance fluxes: A new strategy for high-frequency  
 1570 attenuation correction. *Atmospheric Measurement Techniques*, 12(11), 6059–6078.  
 1571 <https://doi.org/10.5194/amt-12-6059-2019>

- 1572 Mosier, A. R. (2008). Exchange of Gaseous Nitrogen Compounds Between Terrestrial Systems  
 1573 and the Atmosphere. *Nitrogen in the Environment*, 443–462. [https://doi.org/10.1016/B978-](https://doi.org/10.1016/B978-0-12-374347-3.00013-5)  
 1574 [0-12-374347-3.00013-5](https://doi.org/10.1016/B978-0-12-374347-3.00013-5)
- 1575 Mushinski, R. M., Phillips, R. P., Payne, Z. C., Abney, R. B., Jo, I., Fei, S., Pusede, S. E., White,  
 1576 J. R., Rusch, D. B., & Raff, J. D. (2019). Microbial mechanisms and ecosystem flux  
 1577 estimation for aerobic NO<sub>y</sub> emissions from deciduous forest soils. *Proceedings of the*  
 1578 *National Academy of Sciences of the United States of America*, 116(6), 2138–2145.  
 1579 <https://doi.org/10.1073/pnas.1814632116>
- 1580 Neuman, J. A., Trainer, M., Brown, S. S., Min, K. E., Nowak, J. B., Parrish, D. D., Peischl, J.,  
 1581 Pollack, I. B., Roberts, J. M., Ryerson, T. B., & Veres, P. R. (2016). HONO emission and  
 1582 production determined from airborne measurements over the Southeast U.S. *Journal of*  
 1583 *Geophysical Research*, 121(15), 9237–9250. <https://doi.org/10.1002/2016JD025197>
- 1584 Nodeh-Farahani, D., Bentley, J. N., Crilley, L. R., Caputo, C. B., & VandenBoer, T. C. (2021). A  
 1585 boron dipyrromethene (BODIPY) based probe for selective passive sampling of  
 1586 atmospheric nitrous acid (HONO) indoors. *Analyst*, 146(18), 5756–5766.  
 1587 <https://doi.org/10.1039/d1an01089a>
- 1588 Okiti, I., Efakwu, G., Pindus, M., & Kasak, K. (2025). Environmental and biogeochemical  
 1589 drivers of CH<sub>4</sub> and N<sub>2</sub>O flux variability in treatment wetlands. *Ecological Engineering*,  
 1590 219. <https://doi.org/10.1016/j.ecoleng.2025.107705>
- 1591 Oswald, R., Behrendt, T., Ermel, M., Wu, D., Su, H., Cheng, Y., Breuninger, C., Moravek, A.,  
 1592 Mougín, E., Delon, C., Loubet, B., Pommerening-Röser, A., Sörgel, M., Pöschl, U.,  
 1593 Hoffmann, T., Andreae, M. O., Meixner, F. X., & Trebs, I. (2013). HONO emissions from  
 1594 soil bacteria as a major source of atmospheric reactive nitrogen. *Science*, 341(6151), 1233–  
 1595 1235. <https://doi.org/10.1126/science.1242266>
- 1596 Pan, B., Lam, S. K., Mosier, A., Luo, Y., & Chen, D. (2016). Ammonia volatilization from  
 1597 synthetic fertilizers and its mitigation strategies: A global synthesis. *Agriculture,*  
 1598 *Ecosystems and Environment*, 232, 283–289. <https://doi.org/10.1016/j.agee.2016.08.019>
- 1599 Pan, B., Xia, L., Lam, S. K., Wang, E., Zhang, Y., Mosier, A., & Chen, D. (2022). A global  
 1600 synthesis of soil denitrification: Driving factors and mitigation strategies. *Agriculture,*  
 1601 *Ecosystems and Environment*, 327(September 2021), 107850.  
 1602 <https://doi.org/10.1016/j.agee.2021.107850>
- 1603 Pape, L., Ammann, C., Nyfeler-Brunner, A., Spirig, C., Hens, K., & Meixner, F. X. (2009). An  
 1604 automated dynamic chamber system for surface exchange measurement of non-reactive and  
 1605 reactive trace gases of grassland ecosystems. *Biogeosciences*, 6(3), 405–429.  
 1606 <https://doi.org/10.5194/bg-6-405-2009>
- 1607 Paulot, F., Jacob, D. J., Pinder, R. W., Bash, J. O., Travis, K., & Henze, D. K. (2014). Ammonia  
 1608 emissions in the United States, European Union, and China derived by high-resolution  
 1609 inversion of ammonium wet deposition data: Interpretation with a new agricultural  
 1610 emissions inventory (MASAGE\_NH<sub>3</sub>). *Journal of Geophysical Research*, 119(7), 4343–  
 1611 4364. <https://doi.org/10.1002/2013JD021130>
- 1612 Plake, D., Sörgel, M., Stella, P., Held, A., & Trebs, I. (2015). Influence of meteorology and  
 1613 anthropogenic pollution on chemical flux divergence of the NO-NO<sub>2</sub>-O<sub>3</sub> triad above and

1614 within a natural grassland canopy. *Biogeosciences*, 12(4), 945–959.  
 1615 <https://doi.org/10.5194/bg-12-945-2015>

1616 Plake, D., Stella, P., Moravek, A., Mayer, J. C., Ammann, C., Held, A., & Trebs, I. (2015).  
 1617 Comparison of ozone deposition measured with the dynamic chamber and the eddy  
 1618 covariance method. *Agricultural and Forest Meteorology*, 206, 97–112.  
 1619 <https://doi.org/10.1016/j.agrformet.2015.02.014>

1620 Possanzini, M., Febo, A., & Liberti, A. (1983). New design of a high-performance denuder for  
 1621 the sampling of atmospheric pollutants. *Atmospheric Environment (1967)*, 17(12), 2605–  
 1622 2610. [https://doi.org/10.1016/0004-6981\(83\)90089-6](https://doi.org/10.1016/0004-6981(83)90089-6)

1623 Pugliese, G., Ingrisch, J., Meredith, L. K., Pfannerstill, E. Y., Klüpfel, T., Meeran, K. Byron, J.,  
 1624 Purser, G. Gil-Loaiza, J., van Haren, J., Dontsova, K., Kreuzwieser, J., Ladd, S. N., Werner,  
 1625 C. & Williams, J. (2023). Effects of drought and recovery on soil volatile organic  
 1626 compound fluxes in an experimental forest. *Nature Communications*, 14:5064.  
 1627 <https://doi.org/10.1038/s41467-023-40661-8>

1628 Purchase, M. L., Bending, G. D., & Mushinski, R. M. (2023). Spatiotemporal Variations of Soil  
 1629 Reactive Nitrogen Oxide Fluxes across the Anthropogenic Landscape. *Environmental  
 1630 Science and Technology*, 57(43), 16348–16360. <https://doi.org/10.1021/acs.est.3c05849>

1631 Ramazan, K. A., Syomin, D., & Finlayson-Pitts, B. J. (2004). The photochemical production of  
 1632 HONO during the heterogeneous hydrolysis of NO<sub>2</sub>. *Physical Chemistry Chemical Physics*,  
 1633 6(14), 3836–3843. <https://doi.org/10.1039/b402195a>

1634 Reed, C., Brumby, C. A., Crilley, L. R., Kramer, L. J., Bloss, W. J., Seakins, P. W., Lee, J. D., &  
 1635 Carpenter, L. J. (2016). HONO measurement by differential photolysis. *Atmospheric  
 1636 Measurement Techniques*, 9(6), 2483–2495. <https://doi.org/10.5194/amt-9-2483-2016>

1637 Ren, X., Sanders, J. E., Rajendran, A., Weber, R. J., Goldstein, A. H., Pusede, S. E., Browne, E.  
 1638 C., Min, K. E., & Cohen, R. C. (2011). A relaxed eddy accumulation system for measuring  
 1639 vertical fluxes of nitrous acid. *Atmospheric Measurement Techniques*, 4(10), 2093–2103.  
 1640 <https://doi.org/10.5194/amt-4-2093-2011>

1641 Ren, Y., Stieger, B., Spindler, G., Grosselin, B., Mellouki, A., Tuch, T., Wiedensohler, A., &  
 1642 Herrmann, H. (2020). Role of the dew water on the ground surface in HONO distribution: A  
 1643 case measurement in Melpitz. *Atmospheric Chemistry and Physics*, 20(21), 13069–13089.  
 1644 <https://doi.org/10.5194/acp-20-13069-2020>

1645 Richardson, K., Steffen, W., Lucht, W., Bendtsen, J., Cornell, S. E., Donges, J. F., Drüke, M.,  
 1646 Fetzer, I., Bala, G., Von Bloh, W., Feulner, G., Fiedler, S., Gerten, D., Gleeson, T.,  
 1647 Hofmann, M., Huiskamp, W., Kummu, M., Mohan, C., Nogués-Bravo, D., ... Rockström, J.  
 1648 (2023). *Earth beyond six of nine planetary boundaries*. <https://www.science.org>

1649 Scharko, N. K., Schu, U. M. E., Berke, A. E., Banina, L., Peel, H. R., Donaldson, M. A.,  
 1650 Hemmerich, C., White, R., & Ra, J. D. (2015). *Combined Flux Chamber and Genomics  
 1651 Approach Links Nitrous Acid Emissions to Ammonia Oxidizing Bacteria and Archaea in  
 1652 Urban and Agricultural Soil*. <https://doi.org/10.1021/acs.est.5b00838>

1653 Schindlbacher, A., Zechmeister-Boltenstern, S., & Jandl, R. (2009). Carbon losses due to soil  
 1654 warming: Do autotrophic and heterotrophic soil respiration respond equally? *Global  
 1655 Change Biology*, 15(4), 901–913. <https://doi.org/10.1111/j.1365-2486.2008.01757.x>

1656 Schlesinger, W. H. (2020). *Biogeochemistry: An Analysis of Global Change* (4th ed.). Academic  
1657 Press. <https://doi.org/https://doi.org/10.1016/C2018-0-03255-3>

1658 Seinfeld, J. H., & Pandis, S. N. (2006). *Atmospheric chemistry and physics: from air pollution to*  
1659 *climate change*. (2nd edition). John Wiley & Sons.

1660 Shah, S. B., Grabow, G. L., & Westerman, P. W. (2006). *Ammonia adsorption in five types of*  
1661 *flexible tubing materials*. <https://doi.org/10.13031/2013.22253>

1662 Song, Y., Xue, C., Zhang, Y., Liu, P., Bao, F., Li, X., & Mu, Y. (2023). *Measurement Report:*  
1663 *Exchange Fluxes of HONO over Agricultural Fields in the North China Plain. July, 1–38.*  
1664 <https://zenodo.org/record/8115973>

1665 Sörgel, M., Trebs, I., Wu, D., & Held, A. (2015). A comparison of measured HONO uptake and  
1666 release with calculated source strengths in a heterogeneous forest environment. *Atmospheric*  
1667 *Chemistry and Physics*, 15(16), 9237–9251. <https://doi.org/10.5194/acp-15-9237-2015>

1668 Spataro, F., & Ianniello, A. (2014). Sources of atmospheric nitrous acid: State of the science,  
1669 current research needs, and future prospects. *Journal of the Air and Waste Management*  
1670 *Association*, 64(11), 1232–1250. <https://doi.org/10.1080/10962247.2014.952846>

1671 Stepniewski, W., Stepniewska, Z., & Rozej, A. (2015). Gas Exchange in Soils. *Soil*  
1672 *Management: Building a Stable Base for Agriculture*, 117–144.  
1673 <https://doi.org/10.2136/2011.soilmanagement.c8>

1674 Su, H., Cheng, Y., Oswald, R., Behrendt, T., Trebs, I., Meixner, F. X., Andreae, M. O., Cheng,  
1675 P., Zhang, Y., & Pöschl, U. (2011). Soil nitrite as a source of atmospheric HONO and OH  
1676 radicals. *Science*, 333(6049), 1616–1618. <https://doi.org/10.1126/science.1207687>

1677 Tang, K., Qin, M., Duan, J., Fang, W., Meng, F., Liang, S., Xie, P., Liu, J., Liu, W., Xue, C., &  
1678 Mu, Y. (2019). A dual dynamic chamber system based on IBBCEAS for measuring fluxes  
1679 of nitrous acid in agricultural fields in the North China Plain. *Atmospheric Environment*,  
1680 196(May 2018), 10–19. <https://doi.org/10.1016/j.atmosenv.2018.09.059>

1681 Tang, K., Qin, M., Fang, W., Duan, J., Meng, F., Ye, K., Zhang, H., Xie, P., Liu, J., Liu, W.,  
1682 Feng, Y., Huang, Y., & Ni, T. (2020). An automated dynamic chamber system for exchange  
1683 flux measurement of reactive nitrogen oxides (HONO and NOX) in farmland ecosystems of  
1684 the Huaihe River Basin, China. *Science of the Total Environment*, 745(X), 140867.  
1685 <https://doi.org/10.1016/j.scitotenv.2020.140867>

1686 Tian, H., Pan, N., Thompson, R. L., Canadell, J. G., Suntharalingam, P., Regnier, P., Davidson,  
1687 E. A., Prather, M., Ciais, P., Muntean, M., Pan, S., Winiwarter, W., Zaehle, S., Zhou, F.,  
1688 Jackson, R. B., Bange, H. W., Berthet, S., Bian, Z., Bianchi, D., ... Zhu, Q. (2024). Global  
1689 nitrous oxide budget (1980-2020). *Earth System Science Data*, 16(6), 2543–2604.  
1690 <https://doi.org/10.5194/essd-16-2543-2024>

1691 Vaaitinen, O., Metsälä, M., Persijn, S., Vainio, M., & Halonen, L. (2014). Adsorption of  
1692 ammonia on treated stainless steel and polymer surfaces. *Applied Physics B: Lasers and*  
1693 *Optics*, 115(2), 185–196. <https://doi.org/10.1007/s00340-013-5590-3>

1694 VandenBoer, T. C., Brown, S. S., Murphy, J. G., Keene, W. C., Young, C. J., Pszenny, A. A. P.  
1695 P., Kim, S., Warneke, C., De Gouw, J. A., Maben, J. R., Wagner, N. L., Riedel, T. P.,  
1696 Thornton, J. A., Wolfe, D. E., Dubé, W. P., Öztürk, F., Brock, C. A., Grossberg, N., Lefer,  
1697 B., ... Roberts, J. M. (2013). Understanding the role of the ground surface in HONO  
1698 vertical structure: High resolution vertical profiles during NACHTT-11. *Journal of*

1699 *Geophysical Research Atmospheres*, 118(17), 10,155-10,171.  
1700 <https://doi.org/10.1002/jgrd.50721>

1701 VandenBoer, T. C., Young, C. J., Talukdar, R. K., Markovic, M. Z., Brown, S. S., Roberts, J. M.,  
1702 & Murphy, J. G. (2015). Nocturnal loss and daytime source of nitrous acid through reactive  
1703 uptake and displacement. *Nature Geoscience*, 8(1), 55–60.  
1704 <https://doi.org/10.1038/ngeo2298>

1705 Von Der Heyden, L., Wißdorf, W., Kurtenbach, R., & Kleffmann, J. (2022). A relaxed eddy  
1706 accumulation (REA) LOPAP system for flux measurements of nitrous acid (HONO).  
1707 *Atmospheric Measurement Techniques*, 15(6), 1983–2000. [https://doi.org/10.5194/amt-15-](https://doi.org/10.5194/amt-15-1983-2022)  
1708 [1983-2022](https://doi.org/10.5194/amt-15-1983-2022)

1709 Wang, Fu, X., Wu, D., Wang, M., Lu, K., Mu, Y., Liu, Z., Zhang, Y., & Wang, T. (2021).  
1710 Agricultural Fertilization Aggravates Air Pollution by Stimulating Soil Nitrous Acid  
1711 Emissions at High Soil Moisture. *Environmental Science and Technology*, 55(21), 14556–  
1712 14566. <https://doi.org/10.1021/acs.est.1c04134>

1713 Wang, Y., Fu, X., Wang, T., Ma, J., Gao, H., Wang, X., & Pu, W. (2022). Large Contribution of  
1714 Nitrous Acid to Soil-Emitted Reactive Oxidized Nitrogen and Its Effect on Air Quality.  
1715 *Environmental Science and Technology*. <https://doi.org/10.1021/acs.est.2c07793>

1716 Wolff, V., Trebs, I., Ammann, C., & Meixner, F. X. (2010). Atmospheric Measurement  
1717 Techniques Aerodynamic gradient measurements of the NH<sub>3</sub>-HNO<sub>3</sub>-NH<sub>4</sub>NO<sub>3</sub> triad using a  
1718 wet chemical instrument: an analysis of precision requirements and flux errors. In *Atmos.*  
1719 *Meas. Tech* (Vol. 3). [www.atmos-meas-tech.net/3/187/2010/](http://www.atmos-meas-tech.net/3/187/2010/)

1720 Wu, D., Deng, L., Liu, Y., Xi, D., Zou, H., Sha, Z., Pan, Y., Hou, L., & Liu, M. (2020).  
1721 *Comparisons of the effects of different drying methods on soil nitrogen fractions: Insights*  
1722 *into emissions of reactive nitrogen gases (HONO and NO)*.  
1723 <https://doi.org/10.1080/16742834.2020.1733388>

1724 Wu, D., Horn, M. A., Behrendt, T., Müller, S., Li, J., Cole, J. A., Xie, B., Ju, X., Li, G., Ermel,  
1725 M., Oswald, R., Fröhlich-Nowoisky, J., Hoor, P., Hu, C., Liu, M., Andreae, M. O., Pöschl,  
1726 U., Cheng, Y., Su, H., ... Sörgel, M. (2019). Soil HONO emissions at high moisture content  
1727 are driven by microbial nitrate reduction to nitrite: tackling the HONO puzzle. *ISME*  
1728 *Journal*, 13(7), 1688–1699. <https://doi.org/10.1038/s41396-019-0379-y>

1729 Wu, D., Zhang, J., Wang, M., An, J., Wang, R., Haider, H., Xu-Ri, Huang, Y., Zhang, Q., Zhou,  
1730 F., Tian, H., Zhang, X., Deng, L., Pan, Y., Chen, X., Yu, Y., Hu, C., Wang, R., Song, Y., ...  
1731 Liu, M. (2022). Global and Regional Patterns of Soil Nitrous Acid Emissions and Their  
1732 Acceleration of Rural Photochemical Reactions. *Journal of Geophysical Research:*  
1733 *Atmospheres*, 127(6), 1–16. <https://doi.org/10.1029/2021JD036379>

1734 Yang, J. Y., Drury, C. F., Jiang, R., Worth, D. E., Bittman, S., Grant, B. B., & Smith, W. N.  
1735 (2024). Reactive nitrogen losses from Canadian agricultural soils over 36 years. *Ecological*  
1736 *Modelling*, 495. <https://doi.org/10.1016/j.ecolmodel.2024.110809>

1737 Young, C. J., Washenfelder, R. A., Roberts, J. M., Mielke, L. H., Ostho, H. D., Tsai, C.,  
1738 Pikelnaya, O., Stutz, J., Veres, P. R., Cochran, A. K., Vandenboer, T. C., Flynn, J.,  
1739 Grossberg, N., Haman, C. L., Lefer, B., Stark, H., Graus, M., Gouw, J. De, Gilman, J. B., ...  
1740 Brown, S. S. (2012). *Vertically Resolved Measurements of Nighttime Radical Reservoirs in*  
1741 *Los Angeles and Their Contribution to the Urban Radical Budget*.

1742 Zhang, S., Sarwar, G., Xing, J., Chu, B., Xue, C., Sarav, A., Ding, D., Zheng, H., Mu, Y., Duan,  
1743 F., Ma, T., & He, H. (2021). Improving the representation of HONO chemistry in CMAQ  
1744 and examining its impact on haze over China. *Atmospheric Chemistry and Physics*, *21*(20),  
1745 15809–15826. <https://doi.org/10.5194/acp-21-15809-2021>  
1746 Zhou, S., Young, C. J., VandenBoer, T. C., Kowal, S. F., & Kahan, T. F. (2018). Time-Resolved  
1747 Measurements of Nitric Oxide, Nitrogen Dioxide, and Nitrous Acid in an Occupied New  
1748 York Home. *Environmental Science and Technology*, *52*(15), 8355–8364.  
1749 <https://doi.org/10.1021/acs.est.8b01792>  
1750 Zörner, J., Penning De Vries, M., Beirle, S., Sihler, H., Veres, P. R., Williams, J., & Wagner, T.  
1751 (2016). Multi-satellite sensor study on precipitation-induced emission pulses of NO<sub>x</sub> from soils  
1752 in semi-arid ecosystems. *Atmospheric Chemistry and Physics*, *16*(14), 9457–9487.  
1753 <https://doi.org/10.5194/acp-16-9457-2016>

1 Supporting Information for:

2 **Modification and validation of a commercial dynamic chamber for reactive nitrogen and**  
3 **greenhouse gas flux measurements**

4 Moxy Shah<sup>1</sup>, Kifle Z. Aregahegn<sup>1</sup>, Danial Nodeh-Farahani<sup>1</sup>, Leigh R. Crilley<sup>1,†</sup>, Tasnia Hasan<sup>1</sup>,  
5 Yashar Ebrahimi-Iranpour<sup>1</sup>, Fahim Sarker<sup>1</sup>, Nick Nickerson<sup>2</sup>, Chance Creelman<sup>2</sup>, Sarah Ellis<sup>2</sup>,  
6 Alexander Moravek<sup>1,§</sup>, Trevor C. VandenBoer<sup>1,\*</sup>

7

8 <sup>1</sup> Department of Chemistry, York University, Toronto, Ontario, Canada

9 <sup>2</sup> Eosense Inc., Dartmouth, Nova Scotia, Canada

10 <sup>†</sup> Now at: Atmospheric Services, WSP Australia, Brisbane, QLD, Australia

11 <sup>§</sup> Now at: German Environment Agency, Department of Air Quality, Dessau-Rosslau, Germany

12

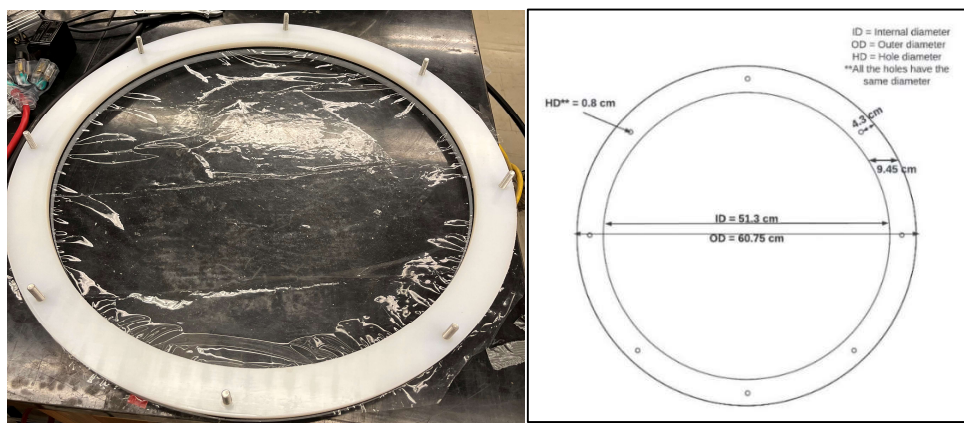
13 \*Communicating author: tvandenb@yorku.ca

14

## 15 S1. Chamber customized components and LabVIEW software

16 To meet the performance requirements for measuring  $N_r$  gases, chamber modifications were  
17 implemented to create inert surfaces that transmit them effectively to downstream gas analyzers.  
18 In addition, a reference chamber with an inert surface at the bottom (Figure S1) was required to  
19 benchmark the performance of a modified chamber (Figure S2) where fittings and surfaces  
20 identified to be reactive or to facilitate stronger surface interactions were replaced with more inert  
21 components.

22



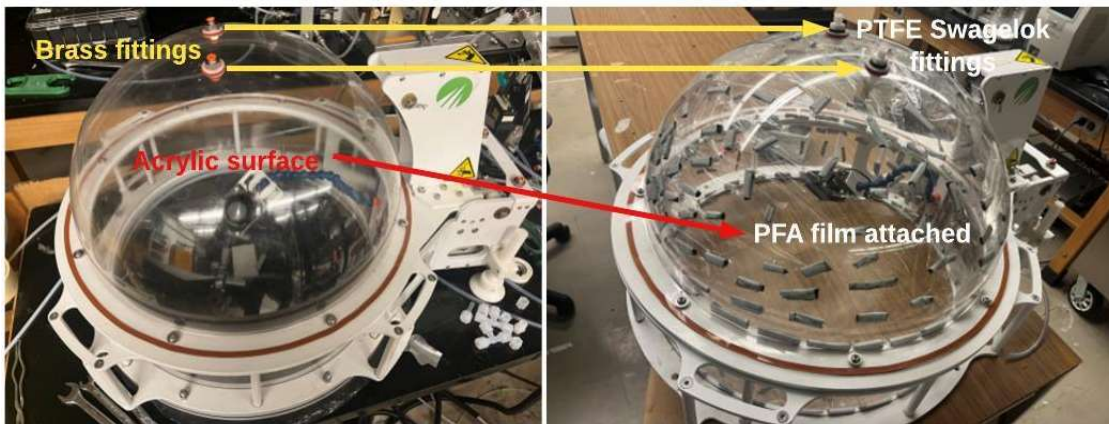
23

24 **Figure S1.** (left) Image of the custom-made base rings made from 1.2 cm thick PTFE with a PFA  
25 film pinned between them; all the holes have identical diameters, are located equidistant from the  
26 edges of the plate and can fit M8 bolts secured with corresponding nuts (McMasterr-Carr®, PN:  
27 90591A161 and 91263A918). (right) Technical scale drawing of a custom-made ring for the  
28 chambers with the inner diameter, outer diameter, ring width, and diameter of holes and their  
29 distance from the outer edges denoted.

30 A custom LabVIEW program was created to automate control of the solenoid valves and mass  
31 flow controller in the measurement set up using the microcontroller. The front panel of the  
32 LabVIEW program (Figure S3a) allows for the timings of the solenoid valve changes to be  
33 manually set, along with the log intervals. [The VI controls valve states \(i.e., for V1, V2, and V3\)](#)  
34 [and synchronization of the valve switching with respect to each chamber opening and closing cycle](#)  
35 [as described in the previous section. The VI also generates its own text file containing valve](#)  
36 [open/close state and a timestamp to be used for data processing. A custom R script \(R Studio](#)  
37 [v3.0.1\) was developed to process the data file generated by eoslink-AC, data from all reactive gas](#)  
38 [analyzers, and the VI.](#) A graphical representation of the valve states in real time is ~~also~~ included  
39 on the front window to allow for the user to assess the measurement state. In the back window  
40 (Figure S3b), the valve timings are controlled within a state machine to allow for continuous valve

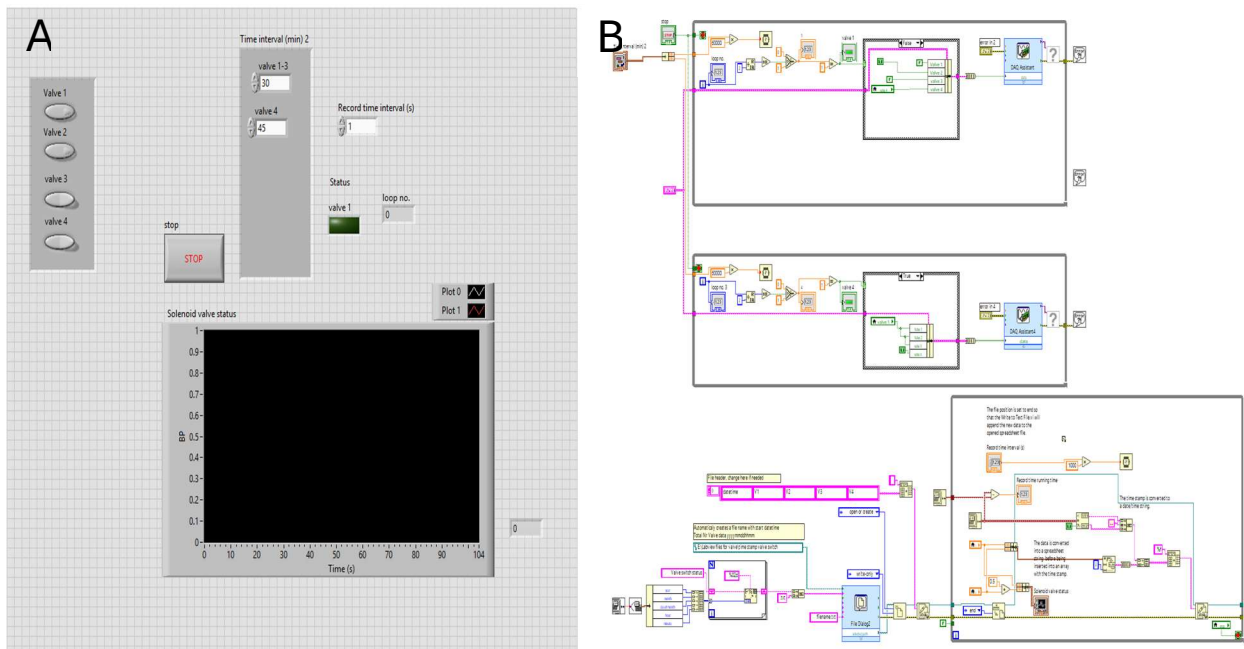
41 switching at the time interval selected on the front window. The MFC flow rate is controlled by  
42 sending the required voltage to the MFC ~~based on the selected flowrate~~. The data logging of the  
43 valve switching timings and MFC flowrates are handled as well within a state machine to ensure  
44 continuous logging to a csv file. The LabVIEW VI is available on the GitHub repository  
45 (<https://github.com/fjs-vdmlab/fluxchamber.git>).

46  
47



48  
49 **Figure S2.** (left) The chamber is in its original configuration from the manufacturer with brass  
50 push-connect inlet and outlet fittings and an acrylic dome. (right) The modified chamber with 1/4”  
51 PTFE Swagelok bulkhead fittings and an attached film of PFA on the interior of the acrylic dome  
52 and all chamber sidewalls using adhesive tape.

53  
54  
55



56  
 57 **Figure S3.** The custom LabVIEW VI for controlling the reference and measurement chambers (A)  
 58 flows, valve states, timing for lid closures, and rate of datalogging through a graphical user  
 59 interface front window, and (B) the back window virtual machines controlling the objects,  
 60 voltages, and data flows.

61

62

63

## 64 S2. Gas analyzer quality control procedures

65 The NO<sub>x</sub> analyzer is calibrated with multipoint NO and NO<sub>2</sub> mixing ratios to generate calibration  
 66 curves on at least an annual basis, with span checks and blanks performed regularly prior to the  
 67 collection of experimental data. For field use, a full calibration is performed prior to relocation of  
 68 the system to the deployment site, a span check on site, and again at the end of the deployment  
 69 period to account for any instrumental drift. Zero air from a gas-calibration instrument (Gascal  
 70 1100TS, American Ecotech, Warren, RI) is introduced to the analyzer through the sampling inlet  
 71 to set the zero point. If the NO, NO<sub>2</sub> and NO<sub>x</sub> readings are within  $\pm 0.5$  ppbv of one another, the  
 72 zero offset is calculated and applied. Then, an NO span check is performed using a known mixing  
 73 ratio from a certified cylinder (e.g. Praxair, NI NO5MC-A3, 4.88 ( $\pm 5\%$ ) ppmv, Toronto, ON) to  
 74 ensure the NO and NO<sub>x</sub> readings are within  $\pm 1$  ppbv of the true value. Next, a multi-point precision  
 75 check of NO is performed at 40 ppbv, 60 ppbv, and 80 ppbv for 30 min, ensuring a stable signal  
 76 has been achieved. The calibration is considered successful when the analyzer's response to

77 changes in NO mixing ratio is linear, with a slope of  $1.00 \pm 0.02$  and  $R^2 > 0.98$ . For multi-point  
78 precision checks of NO<sub>2</sub>, 40 ppbv, 60 ppbv, and 80 ppbv were produced using gas-phase titration  
79 of excess NO with O<sub>3</sub>, again using the GasCal. The criteria for a successful calibration are the  
80 same as described for NO. Conversion efficiency (CE) of NO<sub>2</sub> to NO in the analyzer is the final  
81 parameter to be quantified, and a CE > 96% indicates that the molybdenum catalyst in the analyzer  
82 is operating reliably.

83 Nitrous acid (HONO) was generated using our in-house calibration source based on the reaction  
84 between gas-phase hydrochloric acid (HCl) and a NaNO<sub>2</sub> crystalline film on the surface of PFA  
85 tubing (RS-1) at 50% RH to generate gas-phase HONO and to confirm its unity conversion (Lao  
86 et al., 2020).



88 In the HONO calibration source, a flow of dry carrier gas (Air Ultra Zero, 99.999%, AI 0.0UZ-K,  
89 Praxair) at 50 cm<sup>3</sup> min<sup>-1</sup> passes through a permeation device (PD) containing HCl solution, heated  
90 to 30-40 °C. Another 50 cm<sup>3</sup> min<sup>-1</sup> passes through a glass impinger containing deionized water to  
91 become saturated with water vapour, so the flow obtains an RH of 50% when combined with that  
92 of the HCl. This HCl-water vapour mixture enters a NaNO<sub>2</sub>-coated PFA reaction device, and  
93 HONO is released by acid displacement. A delivered mixing ratio of 10-100 ppb was found to be  
94 effectively converted, in line with the performance we have previously described in detail.

95 Calibration for the O<sub>3</sub> analyzer involves a multipoint check, after one hour of warmup time to  
96 ensure lamp stability, using 100, 200, 300, and 400 ppbv mixing ratios. The procedure is  
97 considered successful if the slope falls between 0.98 and 1.02, the y-intercept between -2 and 2  
98 ppb, and an overall  $R^2 > 0.99$ .

99 The calibration and precision check procedure for the Picarro G2509 involved measuring multiple  
100 known concentrations of GHGs from a calibration cylinder. The calibration cylinder contained a  
101 custom blend of GHGs (4.30 ppm ( $\pm 5\%$ ) N<sub>2</sub>O, 24.8 ppm ( $\pm 5\%$ ) CH<sub>4</sub>, 3587 ppm ( $\pm 2\%$ ) CO<sub>2</sub>),  
102 which was a certified standard grade in an air matrix (Linde Canada Plc; PN: AI CD.4MN1C-A3;  
103 CGA-590). Pure nitrogen and three calibration points at none, 4.5, and 10-times dilution of the  
104 original GHG mixing ratios in the cylinder were delivered to the Picarro analyzer until stable  
105 responses were achieved, in triplicate. The custom-built permeation oven for HONO was also used  
106 to generate known mixing ratios of NH<sub>3</sub> by quantification of bubbler-scrubbed NH<sub>3</sub> using ion

107 chromatography (Salehpoor and VandenBoer, 2023). A dry  $90 \text{ cm}^3 \text{ min}^{-1}$  zero air flow was passed  
108 over a PD containing  $\text{NH}_4\text{OH}$  (30% v/v in 1/8" OD tubing with a 9 cm length). Calibration was  
109 performed for mixing ratios of 30 ppbv to 110 ppbv. The dilution of gases was achieved by  
110 combining the flow from a liquid  $\text{N}_2$  dewar (Linde Canada Plc, PN: NI LC250-230). The observed  
111 versus returned slopes of 0.99 for  $\text{N}_2\text{O}$ , 0.99 for  $\text{CH}_4$ , ~~and~~ 0.98 for  $\text{CO}_2$ , and 1.37 for  $\text{NH}_3$  when  
112 compared to the known concentrations delivered, thereby demonstrating suitable and stable  
113 calibration of the instrument, as these slopes are within the cylinder or IC mixing ratio  
114 uncertainties.

115

116

### 117 **S3. Gas handling for chamber filling and emptying challenge experiments**

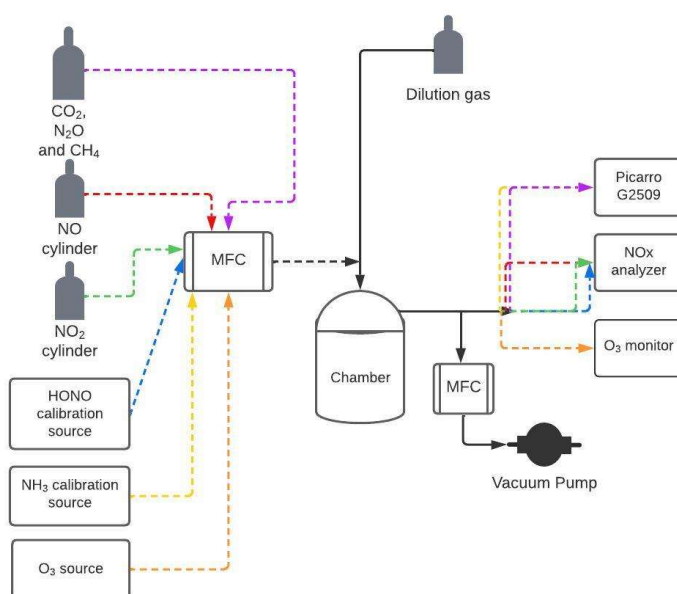
118 Dilution and analyte gas flow rates were controlled using MFCs ( $10 \text{ L min}^{-1}$ , PN:  
119 1179C01314CR1BV, MKS instruments Inc, Andover, MA, US). Mixing ratios of  $\text{CO}_2$ ,  $\text{CH}_4$  and  
120  $\text{N}_2\text{O}$  were controlled by dilution of a calibration gas cylinder containing a custom blend of the  
121 three gases (4000 ppm  $\text{CO}_2$ , 20 ppm  $\text{CH}_4$ , 3.3 ppm  $\text{N}_2\text{O}$  in zero air from Linde Canada). Known  
122 concentrations of  $\text{NO}$  and  $\text{NO}_2$  were added from their respective calibration gas cylinders (5.9 ( $\pm 5$   
123 %) ppmv of  $\text{NO}$  in air, PN: NI NO5MC-A3; 4.5 ( $\pm 5$  %) ppmv  $\text{NO}_2$  in air, PN: NI NX5MC-AQ;  
124 Linde Canada plc, Toronto, ON). Production of known mixing ratios of  $\text{O}_3$  to the chamber was  
125 achieved by combining output from the built-in ozone generator and photometer of the GasCal  
126 into dilution gas at  $2 \text{ L min}^{-1}$ . The  $100 \text{ cm}^3 \text{ min}^{-1}$  flow of HONO exiting the calibration source was  
127 diluted using ZA and mixing ratios spanning 10-100 ppbv were obtained, with  $2 \text{ L min}^{-1}$  directed  
128 to the experimental setup. For  $\text{NH}_3$ , the permeation device output was also diluted with  $2 \text{ L min}^{-1}$   
129 of zero air to generate a mixing ratio on the order of 30 ppbv.

130 The time response of gases for the fill and empty process was calculated from the time required to  
131 fill the chamber and completely empty the chamber (Figure S4). It can be described by a single  
132 exponential function for non-reactive gases such as  $\text{N}_2\text{O}$ ,  $\text{CH}_4$ ,  $\text{CO}_2$ , and  $\text{NO}$  and thus the time  
133 response,  $\tau$ , for the exchange of sample gases volume (Section 2.1; E1-4). The double exponential  
134 function was used for surface-sensitive and reactive gases, such as  $\text{NO}_2$ , HONO, and  $\text{NH}_3$  to yield  
135 two time constants,  $\tau_1$  and  $\tau_2$ , the time response towards the exchange of the sample air volume  
136 and wall interactions (ES-1), respectively (Ellis et al., 2010).

137  $f(t) = y_0 + A_1 \times \exp\left(\frac{-(t-t_0)}{\tau_1}\right) + A_2 \times \exp\left(\frac{-(t-t_0)}{\tau_2}\right)$  ES-1

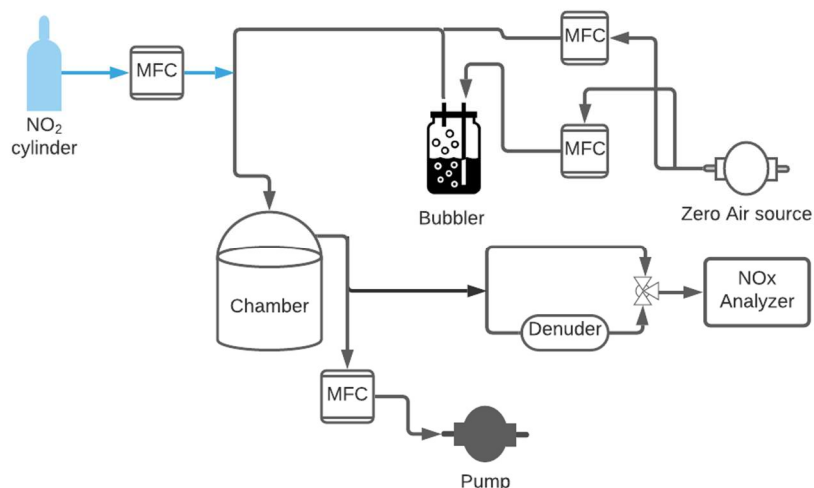
138 Where  $t_0$  is the start time and  $y_0$  is the offset that represents the measurement baseline level;  $A_1$  and  
 139  $A_2$  are proportionality coefficients from the contribution of the physical processes of sample  
 140 volume exchange in the chamber, and reaction and wall interactions, respectively. The relative role  
 141 of wall interactions or reactions,  $D$ , to the overall transfer of gases through a handling system is  
 142 determined by the contribution of the  $A_2$  term to the sum of both  $A$  terms (ES-2).

143  $D = \left(\frac{A_2}{A_1 + A_2}\right) \times 100\%$  ES-2



144  
 145 **Figure S4.** Experimental setup to perform fill/empty experiments in the chambers with greenhouse  
 146 gases: CH<sub>4</sub>, CO<sub>2</sub> and N<sub>2</sub>O (purple); N<sub>r</sub> gases NO (red), NO<sub>2</sub> (green), HONO (blue), NH<sub>3</sub> (yellow),  
 147 and O<sub>3</sub> (orange); and dilution gas (black). Dashed arrows show the flow of gases from their  
 148 respective sources for filling experiments to their respective gas analyzers. Emptying experiments  
 149 use dilution gas alone once a steady state of the challenge gas has been obtained in the chamber.

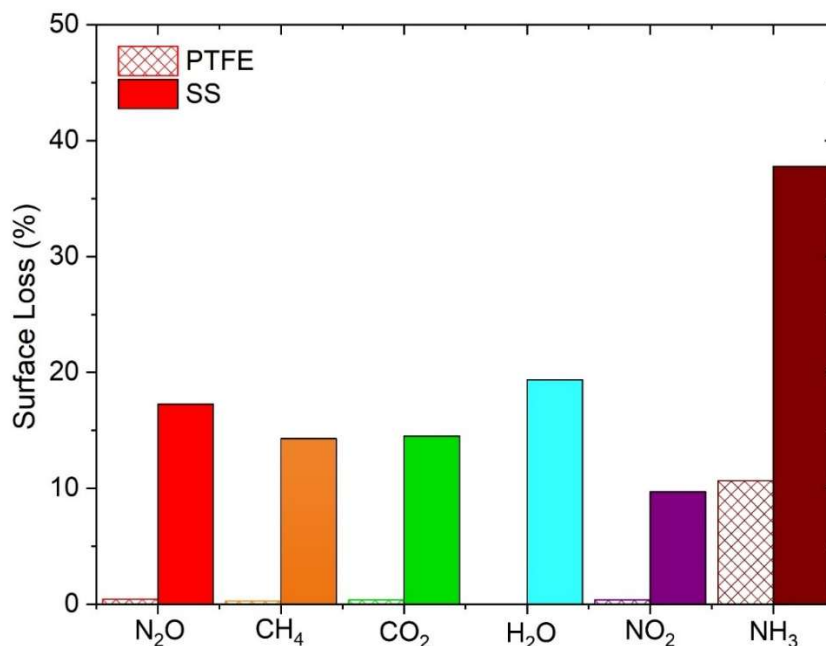
150



151  
 152 **Figure S5.** Experimental setup for NO<sub>2</sub> surface loss characterization, where arrows indicate  
 153 direction of gas flows. Characterization of losses or transformation to HONO was made by a  
 154 modified NO<sub>x</sub> analyzer with experimental flow balance by an added MFC and pump.

155  
 156 **S4. Multiplexer modification tests on gas transfer of N<sub>r</sub> and GHGs**

157 A 2 L min<sup>-1</sup> flow of dry zero air containing 0.058 ppm of N<sub>2</sub>O, 0.57 ppm of CH<sub>4</sub>, and 81.5 ppm of  
 158 CO<sub>2</sub>, was passed through a multiplexer equipped with either PTFE or stainless-steel (SS) valves  
 159 and fittings for 30 min to access relative losses. A separate experiment was performed with 2 L  
 160 min<sup>-1</sup> zero air containing a known concentration of NH<sub>3</sub> (0.41 ppm), which was generated based  
 161 on the method described in Crilley et al. (2023) from a permeation tube containing NH<sub>4</sub>OH.  
 162 Similarly, the transfer efficiency of NO<sub>2</sub> was determined using a calibration cylinder (Linde  
 163 Canada; PN: NI NX5MC-AQ; 5.9 (±5%) ppm) which was combined with a 2 L min<sup>-1</sup> dilution flow  
 164 of dry ZA to achieve a mixing ratio of 54 ppb. All challenge gases were passed through the  
 165 multiplexer equipped with either PTFE or SS valves and PFA or SS fittings, respectively, to assess  
 166 relative losses. All delivered and exiting flows were confirmed to be equal during these tests,  
 167 ensuring no leaks led to a measured loss. The loss percentage in SS fittings and gas handling valves  
 168 was measurable but generally remained below 20%, except for NH<sub>3</sub>, which exhibited a 38%  
 169 reduction. With the modifications, minimal surface loss of GHGs and NO<sub>2</sub> (<1%) was observed  
 170 when using the PFA fittings and PTFE valves, whereas NH<sub>3</sub> exhibited a reduced 11% loss (Figure  
 171 S6).



172

173 **Figure S6.** Loss percentage of GHGs (N<sub>2</sub>O, CH<sub>4</sub>, CO<sub>2</sub>,) and N<sub>r</sub>(NO<sub>2</sub>, NH<sub>3</sub>) gases on stainless steel  
 174 fittings than that of a PTFE valve and unions.

175

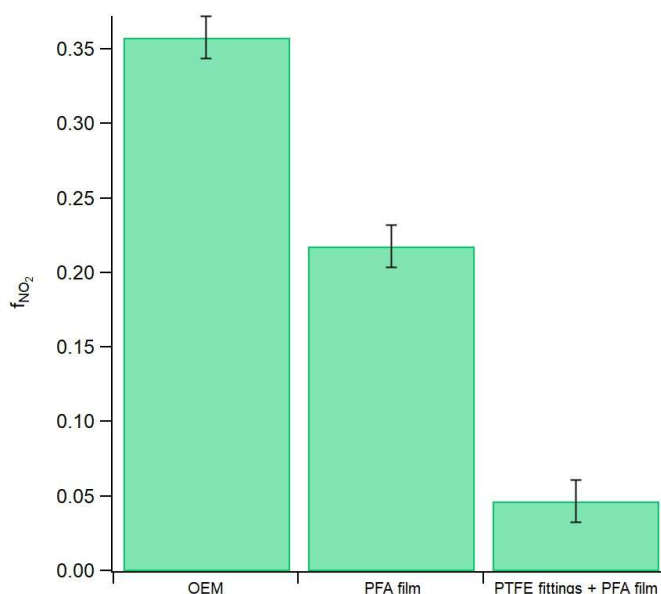
176 The surface porosity of SS can influence its adsorption capacity for CH<sub>4</sub> through van der Waals  
 177 interaction potentials (Sapag et al., 2010). Previous studies have shown that passivation of the SS  
 178 surface via NO treatment can remove surface iron and promote the formation of chromium-  
 179 enriched oxide films, which in turn govern the surface chemical reactivity toward NO (Ma et al.,  
 180 2020). Furthermore, thin layers of iron oxides formed on SS surfaces can affect the interaction  
 181 with N<sub>2</sub>O and alter its decomposition pathways. Consequently, the combined effects of SS surface  
 182 porosity on CH<sub>4</sub> van der Waals interactions, chromium-oxide passive films influencing NO  
 183 reactivity, and iron-oxide layers acting as active adsorption sites for N<sub>2</sub>O may contribute to up to  
 184 a 20 % loss of these gases on SS surfaces. It is likely that the surface losses observed here for very  
 185 clean SS purged extensively with dry zero air prior to these experiments is passivated over time  
 186 when exposed to ambient air continuously, where the surface adsorption sites are fully occupied  
 187 and losses no longer occur.

188

189 **S5. Chamber modification impacts on gas transfer of NO<sub>2</sub> [and materials for O<sub>3</sub>](#)**

190 Gas interactions on chamber surfaces during challenge experiments were reversible for all N<sub>r</sub> and  
 191 GHGs tested, except NO<sub>2</sub>, which showed reactive losses. These were characterized by calculating

192 the lost fraction from a small 5 ppb mixing ratio delivered into an unmodified chamber at a high  
193 relative humidity of 83%, followed by modifications to reduce this outcome under these  
194 challenging simulated environmental conditions (Figure S5). Replacement of the brass push-to-  
195 connect fittings delivered the greatest reduction in reactive losses of NO<sub>2</sub>, with modest gains  
196 obtained from covering chamber surfaces with the PFA film (Figure S7). The final lost fraction of  
197 NO<sub>2</sub> in the modified system reported here is technically at the modified NO<sub>x</sub> analyzer detection  
198 limit and therefore represents a conservative upper limit estimate.



199  
200 **Figure S7.** Average NO<sub>2</sub> loss fraction observed upon addition of 5 ppbv of NO<sub>2</sub> at 83% RH to an  
201 eosAC-LT® chamber (n = 3 each). These trials were conducted in chambers sequentially: without  
202 any modifications (original equipment manufacturer; OEM), following attachment of a PFA film  
203 to the inner surface (PFA film), and replacing brass-lined push to connect fittings with ¼” PTFE  
204 Swagelok bulkhead (PTFE fittings + PFA film).

205  
206 [Different RH values were produced inside the flux chamber by combining flows of dry zero air](#)  
207 [\(ZA\) and one saturated with water vapour by transiting a 500 mL Pyrex impinger filled with](#)  
208 [deionized water \(e.g. equal 1 L min<sup>-1</sup> flows combine to an RH of 50%\). A high-precision, research-](#)  
209 [grade humidity probe \(HMP60, Vaisala Oyj, Finland; ± 3% at 0 - 90% RH, ± 5% at 90 – 100%](#)  
210 [RH\) was connected to the chamber to confirm the set point by measurement in the chamber. The](#)  
211 [flow from an NO<sub>2</sub> calibration cylinder \(Linde Canada; PN: NI NX5MC-AQ; 5.9 \(±5%\) ppm\) was](#)  
212 [adjusted using an MFC and combined with a dilution flow of ZA to achieve mixing ratios of 5, 7](#)  
213 [and 10 ppbv. The total flow rate for these experiments was 2.0 L min<sup>-1</sup> and a vacuum pump \(62.3](#)

214 L min<sup>-1</sup>, PN: UZ-07061-22, Gast Manufacturing Inc., Benton Harbor, MI, USA) with an MFC (10  
215 L min<sup>-1</sup>, PN: 1179C01314CR1BV, MKS instruments Inc, Andover, MA, US) was used to make up  
216 the sampling flow beyond the 0.5 L min<sup>-1</sup> of the NO<sub>x</sub> analyzer. The same flow difference stated  
217 above was maintained in the chamber, using the built-in vent.

218 Prior to O<sub>3</sub> addition, in the characterization experiments, the chamber was flushed with ZA until  
219 the background level of O<sub>3</sub> was at the instrument detection limits. Three mixing ratios of O<sub>3</sub> of  
220 150, 200 and 250 ppb generated by a GasCal 1100 dilution calibrator with integrated O<sub>3</sub> generator  
221 and photometer (American Ecotech, Warren, RI) and added to the chamber for 60 mins or until a  
222 constant concentration was reached. Two replicate runs of each addition level were performed and  
223 delivered mixing ratios were confirmed from bypassing the chambers with just the 15 m gas  
224 handling lines.

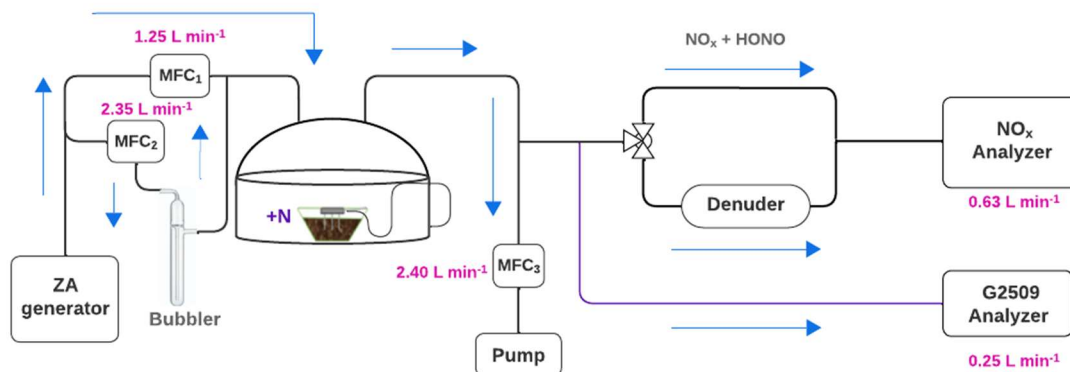
225 Ozone is known to undergo heterogeneous surface reactions, particularly on materials like glass,  
226 metals, or polymers (Plake et al., 2015). These surfaces often contain reactive sites such as  
227 hydroxyl groups or adsorbed water molecules that can catalyze the decomposition of O<sub>3</sub> into  
228 molecular oxygen (O<sub>2</sub>) and other byproducts (George et al., 2015). Adsorption of O<sub>3</sub> on chamber  
229 surfaces is another potential pathway for loss. When O<sub>3</sub> molecules interact with surfaces, they may  
230 undergo a reaction if the materials are not inert. The use of PFA here was intended to benefit the  
231 transfer of O<sub>3</sub>, yet despite its high chemical inertness, low reactivity, and resistance to the uptake  
232 of many chemicals from gas samples, the subsequent O<sub>3</sub>-driven surface reactions are still  
233 substantial (Ebnesajjad, 2017). This is likely because, over time, reactive substances from sampled  
234 air accumulated through adsorption on the PFA film and/or its defect sites, thereby reducing its  
235 effectiveness. The same issue is well-known for PFA tubing used in standard air quality monitoring  
236 and is a common maintenance need for O<sub>3</sub> analyzer inlets. Further, physical wear or surface aging  
237 might alter the material surface through product formation or exposure of new reaction sites, which  
238 makes it less resistant to further reaction with O<sub>3</sub>, and therefore increases the rate of loss over time.  
239 This is the case in our aged PFA film results, which emphasizes the importance of regular  
240 replacement of the film as part of the N<sub>r</sub> system maintenance and that quality control procedures  
241 characterizing the chamber material performance with respect to O<sub>3</sub> will provide the highest  
242 accuracy of experimental results.

243

244 **S6. Agricultural soil samples used for laboratory  $N_r$  emission measurements**

245 A systematic method was used to collect soil samples from an agricultural field to avoid sampling  
246 bias, assuming there is no topographic reason for differences in soil properties and nutrient content  
247 within the specified plot area. The total field area was divided into eight identical plots (A-H) of  
248  $10,800 \text{ m}^2$  ( $120 \text{ m} \times 90 \text{ m}$ ). Soil samples were collected from four random coordinates within each  
249 plot, assuming that samples collected within  $1 \text{ m}^2$  were homogeneous. An approximate 1 L soil  
250 core  $\sim 15 \text{ cm}$  deep was collected with a clean shovel into a pre-labelled clean Ziploc bag. All soil  
251 samples were stored below  $10 \text{ }^\circ\text{C}$  and transported to the laboratory within a week.

252 Four within-plot samples were collected to better approximate soil properties and nutrient content  
253 across the entire field. A total of 32 soil samples were collected from the field, where three of the  
254 randomly collected within-plot samples were combined generate a plot scale sample, and one  
255 sample within each plot was randomly selected to be combined into a pooled sample representative  
256 of the field scale. For the purposes of this work, we assessed the emissions of NO, NO<sub>2</sub>, and HONO  
257 from duplicate pooled samples and from Plot D within this sampling experiment (Figure S8).



258

259 **Figure S8.** Schematic representation of a dynamic chamber system to measure NO, NO<sub>2</sub>, and  
260 HONO fluxes with a modified NO<sub>x</sub> analyzer (black lines) and NH<sub>3</sub> and N<sub>2</sub>O for soil amended with  
261 fertilizer using G2509 (purple line). Zero air and humid airflow rates were set using MFC 1 and  
262 MFC 2, for a total of  $3.6 \text{ L min}^{-1}$  going into the chamber. The three-way valve either sends  
263 headspace air directly to the NO<sub>x</sub> analyzer (measures NO and NO<sub>2</sub>+HONO), or via the Na<sub>2</sub>CO<sub>3</sub>-  
264 coated denuder (measures NO, and NO<sub>2</sub>). Blue arrows indicate the flows throughout the  
265 experimental setup, with flow rates provided in pink text.

266

267 Soil NH<sub>3</sub> and N<sub>2</sub>O, along with NO, NO<sub>2</sub>, and HONO, were measured when the pooled soil sample  
268 was amended with four different nitrogen salts (i.e., AN, urea, ABC, AC; Table S1, Figure 4). For

269 each experiment, 350 g of soil was used, and a corresponding fertilizer solution equivalent to 100  
270 kg N ha<sup>-1</sup> was added. Additional experimental details are provided in Section 2.65.1 of the main  
271 manuscript. The interpretation of the full set of experimental results will be presented in a  
272 forthcoming publication on field-scale heterogeneity of N<sub>r</sub> emissions and comments on the  
273 assumption that individual samples represent these processes across fields with similar  
274 physicochemical properties.

275

**Table S1.** The average and integrated flux for soil treated with 4 different nitrogen salts (urea, ammonium bicarbonate (ABC), ammonium carbonate (AC), and ammonium nitrate (AN)). Average fluxes were calculated over the entire experimental cycle. Integrated fluxes were obtained using the trapezoidal rule. Standard deviations reflect the variability in fluxes measured across the time points for the single sample analyzed per treatment. Replicates of urea (~40 hours) and AN (~5 hours) were analyzed with only the modified NO<sub>x</sub> analyzer, thus lacking data for N<sub>2</sub>O and NH<sub>3</sub> (represented by -).

Soil sample	Avg Flux ( $\mu\text{g N m}^{-2} \text{hr}^{-1}$ )					Integrated flux ( $\mu\text{g N m}^{-2}$ )				
	NO	NO <sub>2</sub>	HONO	N <sub>2</sub> O	NH <sub>3</sub>	NO	NO <sub>2</sub>	HONO	N <sub>2</sub> O	NH <sub>3</sub>
Urea	0.1 ± 0.06	0.03 ± 0.01	0.002 ± 0.004	9 ± 3	1 ± 0.5	230	120	1	32400	3700
ABC	0.03 ± 0.01	0.1 ± 0.01	0.03 ± 0.02	7 ± 6	20 ± 7	90	300	90	25300	85900
AC	0.03 ± 0.004	0.2 ± 0.03	0.01 ± 0.01	3 ± 2	20 ± 3	90	500	1	9600	63700
*AN	1 ± 0.1	1 ± 0.1	1 ± 0.02	-	-	50	50	40	-	-

## 1 S7. Analytical solutions to soil reactive gas flux determination with a two-chamber setup

2 For a single chamber system, ES1 can be used to calculate the flux of gases to and from an enclosed  
3 surface. Each of these terms and their units are provided in Section 2.76 of the main manuscript.

$$4 \quad V \frac{dc_{\text{cham}}(t)}{dt} = (A \cdot F_{\text{soil}}(t)) + (Q_{\text{in}} \cdot c_{\text{in}}(t)) - (Q_{\text{out}} \cdot c_{\text{out}}(t)) + (V \cdot R(t)) \quad \text{ES1}$$

5 Where  $V$  is the volume of the chamber ( $\text{m}^3$ ),  $A$  is the surface area ( $\text{m}^2$ ) enclosed by the chamber  
6 and governing the gas flux;  $Q_{\text{out}}$  is the volumetric flow rate of air exiting the chamber ( $\text{m}^3 \text{s}^{-1}$ ;  $c_{\text{m}}(t)$   
7 and  $c_{\text{r}}(t)$  are target gas concentrations within the measurement and reference chamber ( $\text{mol m}^{-3}$ ),  
8 respectively;  $c_{\text{in}}$  and  $c_{\text{out}}$  represent their corresponding rates of change ( $\text{mol m}^{-3} \text{s}^{-1}$ ); and  $F_{\text{net}}$  is the  
9 resulting net gas flux per unit area ( $\text{mol m}^{-2} \text{s}^{-1}$ ). The terms  $R_{\text{r}}$  denote the instantaneous chemical  
10 production or loss rate expressed in units of  $\text{mol m}^{-3} \text{s}^{-1}$  for consistency.

11 In our field pilot study, the measurement and reference chambers were installed on clay soil using  
12 serrated collars inserted about 10 cm deep into pre-dug circles (30 cm diameter, 20 cm depth)  
13 between crop rows of senescing soybeans after removing debris. Collars were left to settle for over  
14 24 h before the chambers were bolted securely to ensure a gas-tight seal. Flux observations were  
15 made for two weeks from 02-16 September 2022, during a period when daytime air temperatures  
16 were regularly over 25 °C and there were no instances of rain. During the second week of  
17 observations, a fertilizer addition of 25 kg N ha<sup>-1</sup> was made by broadcasting urea in the chamber,  
18 followed by applying clean water to simulate 2.5 cm (1”) of rainfall.

19 In our sampling setup, there is a negligible concentration of any target gas in the inflow ( $c_{\text{in}}(t) \approx 0$ )  
20 due to the use of ultra-pure zero air or nitrogen, and the air is actively mixed inside the chamber  
21 with its fan, such that  $c_{\text{chamb}} = c_{\text{out}}$ , so the equation simplifies to ES2. For non-steady-state  
22 conditions, where the concentration inside the chamber is changing over time, the time-varying  
23 nature of the sources and sinks needs to be accounted for. The instantaneous equation that describes  
24 the behaviour over an interval  $[t_1, t_2]$  is the integral in ES3.

$$25 \quad V \frac{dc_{\text{cham}}(t)}{dt} = (A \cdot F_{\text{soil}}(t)) - (Q_{\text{out}} \cdot c_{\text{cham}}(t)) + (V \cdot R(t)) \quad \text{ES2}$$

$$26 \quad \int_{t_1}^{t_2} V \frac{dc_{\text{cham}}(t)}{dt} dt = \int_{t_1}^{t_2} [(A \cdot F_{\text{soil}}(t)) - (Q_{\text{out}} \cdot c_{\text{cham}}(t)) + (V \cdot R(t))] dt \quad \text{ES3}$$

27 Assuming the soil flux,  $F_{\text{soil}}(t)$  ( $\text{mol m}^{-2} \text{s}^{-1}$ ), is constant over the time interval  $[t_1, t_2]$  is a reasonable  
 28 assumption for experimental setups where the environmental conditions and, where relevant, soil-  
 29 plant activity, do not change substantially over short periods of a few to tens of minutes (Pape et  
 30 al., 2009). Rearranging the ES3 expression, a flux can be calculated using ES4.

$$31 \quad F_{\text{soil}} = \frac{V\Delta C_{\text{cham}} + Q_{\text{out}} \int_{t_1}^{t_2} c_{\text{cham}}(t) dt - V \int_{t_1}^{t_2} R(t) dt}{A(t_2 - t_1)} \quad \text{ES4}$$

32 The volumetric mixing ratio (ES5) is often the unit of measurement for many gas analyzers and  
 33 defined as the number of moles of the gas divided by the number of moles of dry air. The mixing  
 34 ratio  $X$  is dimensionless and provides a way to express gas concentrations independently of  
 35 changes in temperature or pressure.

$$X = \frac{n_{\text{gas}}}{n_{\text{air}}} \quad \text{ES5}$$

36 We can then write the molar concentration,  $c$  ( $\text{mol m}^{-3}$ ), in the form of volumetric mixing ratio  
 37 using ES6:

$$c_{\text{chamber}} = X_{\text{out}} \times \rho_d \quad \text{ES6}$$

38 Here,  $\rho_d$  ( $\text{mol m}^{-3}$ ) is the molar density of the air. The total pressure  $P$  is equal to atmospheric  
 39 pressure, such that by using the ideal gas law, we can define the density of air using ES7:

$$\rho_d = \frac{P}{R \cdot T} \quad \text{ES7}$$

40 Where,  $P$  (Pa) is the pressure,  $R$  ( $\text{J mol}^{-1} \text{K}^{-1}$ ) is the universal gas constant, and  $T$  (K) is the absolute  
 41 temperature. In ES7 it is assumed that atmospheric air behaves as an ideal gas, which is a  
 42 reasonable approximation under typical environmental conditions, even at the Earth's surface. The  
 43 resulting molar concentrations can then be calculated and used in the flux determinations as  
 44 presented in E6 and E7 in the main manuscript.

## 45 **S7.1 Dynamics and physical corrections from reference and measurement chambers**

46 In environmental and experimental settings where gas flux is measured, it is important to capture  
 47 the rate at which gas concentrations change over time. This rate reflects the net outcome from both

48 gas emission and deposition occurring simultaneously, influenced by the surface under study,  
 49 temperature, pressure, flow, and surface interactions (ES8). The RC (r), which operates under  
 50 identical conditions to the MC (m) but excludes the surface under study by isolating it beneath the  
 51 PFA film, takes into account the environmental and surface effects, such that its observed flux rate  
 52 is described by (ES9). To isolate the effects of the surface under study from reactions taking place  
 53 in the atmospheric sample and on the chamber surfaces, the rate of change in the RC is subtracted  
 54 from that in the MC, yielding the net flux through ES10.

$$F_m = \text{surface under study} + \text{environmental and chamber effects} \quad \text{ES8}$$

$$F_r = \text{environmental and chamber effects} \quad \text{ES9}$$

$$F_{\text{net}} = F_m - F_r \quad \text{ES10}$$

55 The net rate of change can now be used to calculate the molar flux by substituting it into ES4 to  
 56 arrive at ES11, which is the core of E6.

$$F_{\text{net}} = \frac{V}{A} \left( \frac{\Delta C_m}{\Delta t_m} - \frac{\Delta C_r}{\Delta t_r} \right) + \frac{Q_{\text{out}}}{A} \left( \frac{\int_{t_{1m}}^{t_{2m}} c_m(t) dt}{\Delta t_m} - \frac{\int_{t_{1r}}^{t_{2r}} c_r(t) dt}{\Delta t_r} \right) \quad \text{ES11}$$

$$- \frac{V}{A} \left( \frac{\int_{t_{1m}}^{t_{2m}} R_m(t) dt}{\Delta t_m} - \frac{\int_{t_{1r}}^{t_{2r}} R_r(t) dt}{\Delta t_r} \right)$$

57 The closure period of the two chambers is matched to facilitate an accurate correction of the  
 58 kinetics and surface interactions. The calculated flux from the surface under study is therefore  
 59 corrected for background and environmental effects, reducing systematic bias from gas reactivity.

60 Last, we introduce a flux [attenuation](#) factor ( $\lambda$ ; ES12) to correct for partial transmission of a  
 61 targeted reactive gas due to adsorption to chamber walls and gas handling lines during a  
 62 measurement cycle ( $t_1$ - $t_2$ ). This factor needs to be determined empirically through calibration  
 63 experiments, like those presented in Figure 2, where the loss [of the surface-active gas](#) is quantified  
 64 relative to [the transmission of an inert gas](#) during both filling and emptying processes.

$$\lambda = \frac{\int_{t_1}^{t_2} \left(\frac{dX}{dt}\right)_{\text{theoretical}} dt}{\int_{t_1}^{t_2} \left(\frac{dX}{dt}\right)_{\text{measured}} dt}$$

ES12

65 We can account for  $\lambda$  in ES11, to get the molar flux of gas inside the chamber, which accounts for  
 66 flux loss (ES13) by scaling the equation directly to arrive at the full form presented as E6 in the  
 67 main manuscript.

$$F_{\text{net}} = (\lambda) \cdot \left( \frac{V}{A} \left( \frac{\Delta C_m}{\Delta t_m} - \frac{\Delta C_r}{\Delta t_r} \right) + \frac{Q_{\text{out}}}{A} \left( \frac{\int_{t_{1m}}^{t_{2m}} c_m(t) dt}{\Delta t_m} - \frac{\int_{t_{1r}}^{t_{2r}} c_r(t) dt}{\Delta t_r} \right) \right. \\ \left. - \frac{V}{A} \left( \frac{\int_{t_{1m}}^{t_{2m}} R_m(t) dt}{\Delta t_m} - \frac{\int_{t_{1r}}^{t_{2r}} R_r(t) dt}{\Delta t_r} \right) \right)$$

ES13

68

## 69 **S7.2 Chemical corrections from reference and measurement chambers**

70 For reactive species impacting our  $N_r$  suite – mainly NO, O<sub>3</sub>, and NO<sub>2</sub> – flux values can be  
 71 attenuated or enhanced due to the reaction of NO with O<sub>3</sub> (RS-2). We show below (Figure S9) that  
 72 the photolysis of NO<sub>2</sub> and HONO are negligible due to the high-energy photon cut-off of acrylic  
 73 (<400 nm threshold) (RS-3). Nevertheless, we include the NO<sub>2</sub> photolysis term in the rate  
 74 expression for completeness, noting that under our field conditions its contribution (Section  
 75 S7.2.1) is within the noise of the NO<sub>2</sub> fluxes. The rate of RS-2 and RS-3 for NO<sub>2</sub>, therefore, can  
 76 be expressed by ES14.



79  $\frac{d[\text{NO}_2]}{dt} = k_{\text{NO}+\text{O}_3}[\text{NO}][\text{O}_3] - J_{\text{NO}_2}[\text{NO}_2]$  **ES14**

80 The total change in the concentration due to reaction over a cycle can be expressed by integrating  
 81 the rate equation over the cycle duration, as the start time ( $t_0$ ) and end time ( $t$ ) are known, therefore  
 82 replacing the term  $R$  in ES13 (and E6) with the expression in ES15. Under our chamber conditions,

83 the integrated photolysis contribution was smaller than the measurement uncertainty of the NO<sub>2</sub>  
 84 fluxes (Section S7.2.1), and therefore it was not included as an explicit term in ES15.

$$\int_{t_0}^t R dt = \Delta[\text{NO}] = \Delta[\text{O}_3] = -\Delta[\text{NO}_2] = - \int_{t_0}^t k [\text{NO}(t)][\text{O}_3(t)] dt \quad \text{ES15}$$

85 Where the  $R$  is the loss rate due to the reaction for NO and O<sub>3</sub> ( $R \leq 0$ ) and simultaneous production  
 86 rate for NO<sub>2</sub> ( $R \geq 0$ ). For the kinetic determination, number density of the gases (molec cm<sup>-3</sup>) are  
 87 used to track the chemical transformations over time, but instrumentation typically measures  
 88 mixing ratios. If temperature and pressure are measured alongside target gases in the dynamic  
 89 chambers, then using ES6 and ES7, the reaction-corrected flux based on measured mixing ratios  
 90 can be calculated using ES16.

$$F_{\text{net}} = \lambda \cdot \frac{P_{\text{air}}}{R \cdot T} \cdot \left( \frac{V}{A} \left( \frac{\Delta X_m}{\Delta t_m} - \frac{\Delta X_r}{\Delta t_r} \right) + \frac{Q_{\text{out}}}{A} \left( \frac{\int_{t_{1m}}^{t_{2m}} X_m(t) dt}{\Delta t_m} - \frac{\int_{t_{1r}}^{t_{2r}} X_r(t) dt}{\Delta t_r} \right) \right. \\ \left. - \frac{V}{A} \left( \frac{P_{\text{air}}}{R \cdot T} \right) \left( \frac{\int_{t_{1m}}^{t_{2m}} k \cdot X_{\text{NO}_m}(t) \cdot X_{\text{O}_3_m}(t) dt}{\Delta t_m} - \frac{\int_{t_{1r}}^{t_{2r}} k \cdot X_{\text{NO}_r}(t) \cdot X_{\text{O}_3_r}(t) dt}{\Delta t_r} \right) \right) \quad \text{ES16}$$

93  
 94 Where, again,  $P_{\text{air}}$  (Pa) is the atmospheric pressure inside the measurement and reference  
 95 chambers,  $R$  (J mol<sup>-1</sup> K<sup>-1</sup>) is the gas constant, and  $T$  (K) is the temperature inside the RC and MC.  
 96 For non-reactive gases like CO<sub>2</sub>, CH<sub>4</sub>, and N<sub>2</sub>O, there is no reaction term and the flux can be  
 97 described simply using ES17 as below, which is presented as E7 in the main manuscript.

$$F_{\text{net}} = \lambda \cdot \frac{P_{\text{air}}}{R \cdot T} \cdot \left( \frac{V}{A} \left( \frac{\Delta X_m}{\Delta t_m} - \frac{\Delta X_r}{\Delta t_r} \right) + \frac{Q_{\text{out}}}{A} \left( \frac{\int_{t_{1m}}^{t_{2m}} X_m(t) dt}{\Delta t_m} - \frac{\int_{t_{1r}}^{t_{2r}} X_r(t) dt}{\Delta t_r} \right) \right) \quad \text{ES17}$$

99

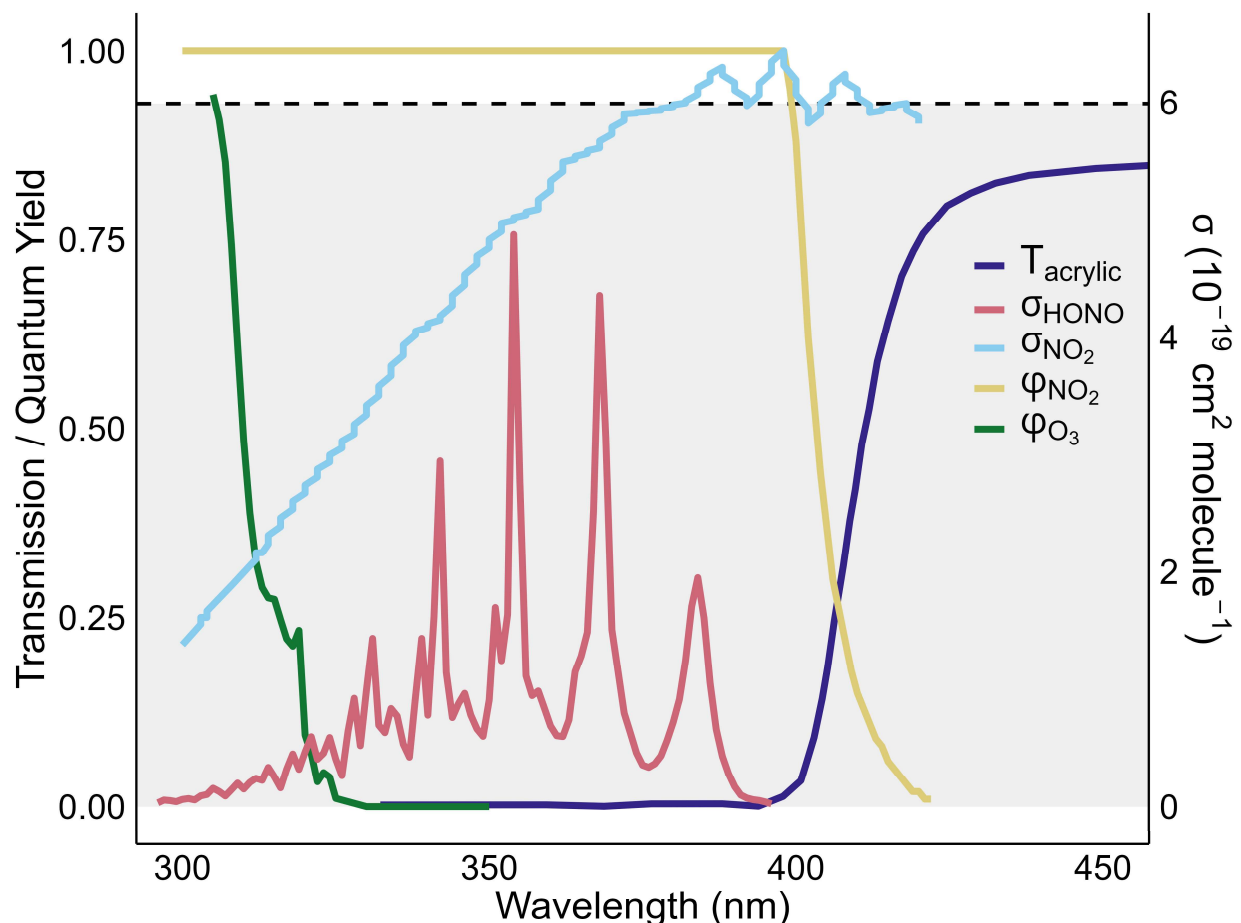
### 100 **S7.2.1 Losses of NO<sub>2</sub> or HONO due to photolysis within transparent chambers**

101 The fraction of NO<sub>2</sub> or HONO lost due to photolysis depends on the photons transmitted through  
 102 the chamber material (Figure S9). The photolysis rate is estimated here under a worst-case scenario  
 103 with respect to our field observations, using actinic flux data collected for Lambton County,

104 Ontario, Canada, at 12:00 PM local time on September 10, 2022, under clear sky conditions. The  
105 effective photolysis rate constant within the chamber was calculated to be  $2.62 \times 10^{-4} \text{ s}^{-1}$ , using  
106 ES18.

$$J = \int [\sigma(\lambda) \cdot \phi(\lambda) \cdot F(\lambda) \cdot T(\lambda)] d\lambda \quad \text{ES18}$$

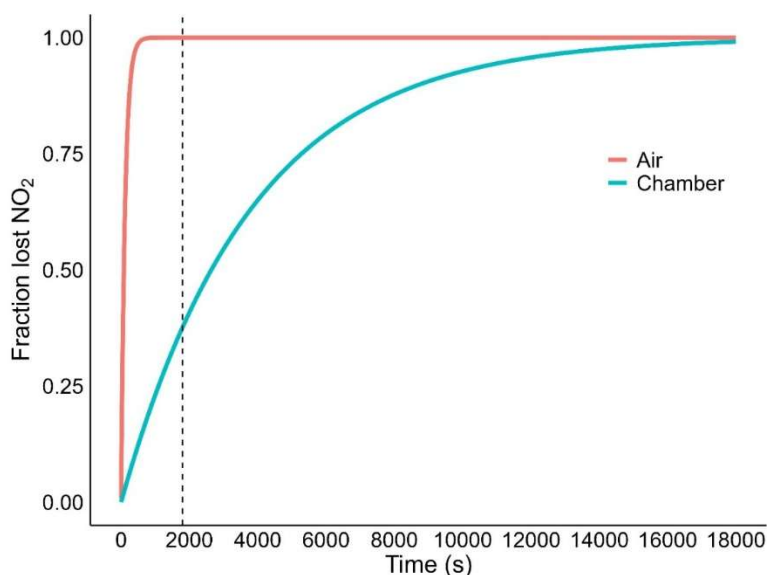
107 Where  $\sigma(\lambda)$  is the absorption cross-section of  $\text{NO}_2$  ( $\text{cm}^2 \text{ molecule}^{-1}$ ),  $\Phi(\lambda)$  is the quantum yield  
108 (dimensionless; 0-1),  $F(\lambda)$  is the actinic flux ( $\text{photons cm}^{-2} \text{ s}^{-1} \text{ nm}^{-1}$ ), and  $T(\lambda)$  is the transmission  
109 of the acrylic chamber material (dimensionless) at wavelength  $\lambda$  (dimensionless; 0-1).



110  
111 **Figure S9.** Transmission of light through the chamber and quantum yields for photolysis of  $\text{O}_3$   
112 and  $\text{NO}_2$ , along with the absorption cross section of  $\text{NO}_2$  and HONO, as a function of wavelength.  
113 The shaded region denotes the manufacturer-specified minimum transmission ( $\geq 93\%$ ) for the PFA  
114 film. Note the negligible overlap for photodissociation. The acrylic effectively blocks wavelengths  
115 below approximately 400 nm, meaning the photolysis rate of  $\text{NO}_2$  is the only one impacted by a  
116 chamber constructed from this material.

117

118 Using this attenuated loss rate in a kinetic model, we find that at most 14% of the initial NO<sub>2</sub>  
119 concentration is lost due to photolysis within the chamber in the first 10 minutes, and 37% within  
120 30 minutes (Figure S10, blue trace). This rate of photolysis significantly reduces the NO<sub>2</sub> mixing  
121 ratio within the chamber environment. The wavelength cut-off of polyacrylate limits the  
122 penetration of shorter wavelengths, which are more effective in driving NO<sub>2</sub> photolysis, so the  
123 overall rate is slowed inside the chamber compared to ambient conditions (Figure S10, red trace).  
124 Despite this, within the initial 30-minute timeframe (Figure S10, vertical dashed line), the change  
125 in the fraction of NO<sub>2</sub> lost is relatively small due to the dynamic nature of the chamber headspace  
126 (i.e. due to continuous sampling and dilution flows), allowing for a reasonable linear  
127 approximation of the curve. The photolytic loss of HONO is much more impacted by the cutoff  
128 wavelength of the polyacrylate chamber lid, to the point where it is not substantially lost due to  
129 photolysis. As shown in Figure S9, the HONO absorption cross-section lies almost entirely below  
130 the ~400 nm transmission cutoff of the chamber lid. This spectral mismatch means that HONO  
131 photolysis is effectively suppressed inside the chamber, unlike NO<sub>2</sub>, which still absorbs within the  
132 transmitted range. Consequently, chamber-derived HONO fluxes are not biased by photolytic loss,  
133 making their interpretation simpler than NO<sub>2</sub>.



134  
135 **Figure S10.** Fraction of NO<sub>2</sub> lost due to photolysis over time, both inside the chamber (blue line)  
136 and in ambient air (red line). The vertical dashed line indicates the 30-min mark, where an initial  
137 linear approximation of the loss rate can be applied.

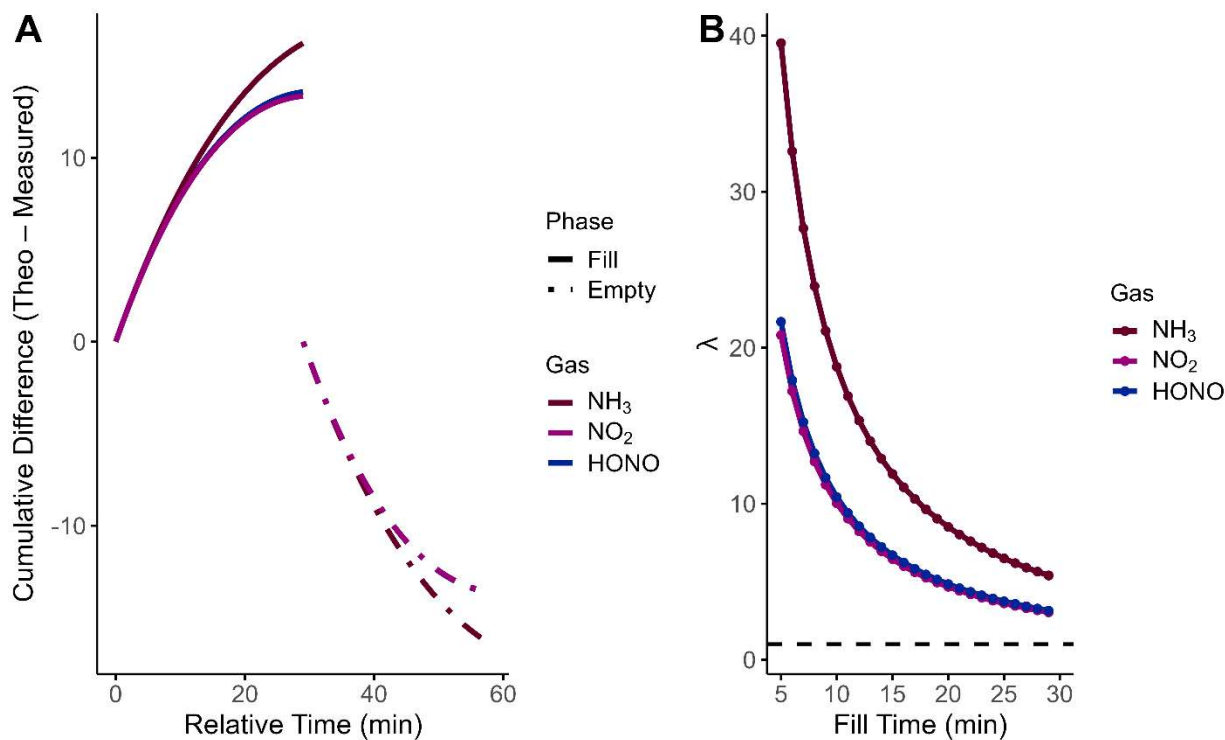
138 This simplification yields an estimated initial photolysis rate loss of approximately 1% per minute  
139 for NO<sub>2</sub> over the initial 30-minute measurement cycle and effectively 0% per minute for HONO

140 under chamber conditions due to the cutoff wavelength. This linear approximation is only valid  
141 for this kind of short initial period. Beyond 30 minutes, the overall trend exhibits a steeper  
142 exponential decay, characteristic of the NO<sub>2</sub> photolysis process. This photolysis contribution is  
143 included within the reaction term (ES14). Its effect (~1% min<sup>-1</sup> maximum at noon under clear-sky  
144 conditions) is smaller than the uncertainties of the NO<sub>2</sub> flux measurements, so it is not explicitly  
145 propagated into the flux equations in this work.

### 146 **S7.3 Surface interaction correction using the attenuation factor ( $\lambda$ )**

147 Accurate quantification of trace gas fluxes using dynamic chamber systems also requires  
148 correction for attenuation caused by gas-wall interactions and the associated sensor response  
149 delays. These effects can significantly suppress the measured accumulation rate of reactive species  
150 within the closed chamber, particularly for compounds with strong surface affinity, such as NH<sub>3</sub>,  
151 NO<sub>2</sub>, and HONO, requiring the use of the attenuation factor ( $\lambda$ ; ES12).

152 The attenuation factor is derived from our controlled delivery experiments that simulated two  
153 sequential measurement scenarios: a filling phase, where gas is introduced and accumulates within  
154 the chamber, and an emptying phase, where the chamber is flushed and the gas signal declines  
155 (Figure 4). For an ideal gas without wall interactions, the integrated difference between theoretical  
156 and observed concentrations within a single chamber should approach zero, and  $\lambda$  would be  
157 approximately one. However, as shown in Figure S11, all three gases exhibited surface  
158 interactions, with NH<sub>3</sub> showing the largest deviation from an inert gas model.



159  
 160 **Figure S11. (A)** Cumulative integration of the difference between inert and measured gas  
 161 concentrations for NH<sub>3</sub>, NO<sub>2</sub>, and HONO over a 30-minute filling and emptying period **and (B)**  
 162 **resulting  $\lambda$  values for modified chambers, 15 m PFA lines, and gas analyzers as a function of filling**  
 163 **time with a dashed line at one corresponding to the transmission of an inert gas.** Theoretical  
 164 concentrations represent expected values from an inert gas without attenuation, while measured  
 165 values reflect signal loss due to gas-wall interactions. Solid lines correspond to the filling phase;  
 166 dashed lines represent the emptying phase, together representing a full in situ measurement cycle.  
 167 These integrals illustrate the time-dependent and nonlinear attenuation that underlies the empirical  
 168 calculation of  $\lambda$ . Note that due to observations switching to the RC, the emptying phase does not  
 169 impact the value of  $\lambda$  applied during field use.

170  
 171 When filling, the gas accumulates in the chamber, with strong initial attenuation that diminishes  
 172 over time as the surfaces passivate. Conversely, adsorbed gases from surface passivation may  
 173 partially desorb during emptying, offsetting some of the losses observed during filling. Two  
 174 respective  $\lambda$  values can be derived and considered separately by integrating the filling and  
 175 emptying concentration-time series independently.

176 Where a single-chamber is used to study an experimental system, the filling- and emptying-phase  
 177 effects offset one another, masking the attenuation as all the surface interacting molecules are  
 178 eventually transferred to the gas analyzer. In field applications with a dual chamber approach in

179 use, this symmetry is lost because the measurement chamber accumulates a real emission flux  
 180 during a defined closure period, while the reference chamber does not. Hence, an empirical  
 181 calculation of  $\lambda$  is needed which compensates for the attenuation in both the accumulation and  
 182 depletion of the target gas (Table S7). For  $\text{NH}_3$ , the fill-phase difference was 16.23, dividing the  
 183 expected theoretical accumulation by this measured accumulation yields a  $\lambda_{\text{fill}}$  of 5.40, while the  
 184 empty-phase difference was  $-16.44$ , with a  $\lambda_{\text{empty}}$  of 0.36. Unsurprisingly,  $\text{NO}_2$  and HONO  
 185 followed a similar trend as we also noted their surface interactions, with  $\lambda_{\text{fill}}$  values of 3.04 and  
 186 3.14 resulting from our characterizations, respectively, and  $\lambda_{\text{empty}}$  values near 0.40. Uncertainties  
 187 in  $\lambda$  were quantified from three replicate fill-empty characterizations under identical laboratory  
 188 conditions, yielding one-sigma deviations of  $\pm 0.21$  for  $\text{NH}_3$ ,  $\pm 0.02$  for  $\text{NO}_2$ , and  $\pm 0.02$  for HONO.  
 189 These values therefore represent attenuation behavior for a pristine or lightly used chamber system.  
 190 From a physical and mathematical standpoint, attenuation reduces all concentration-dependent  
 191 components of the flux (accumulation/depletion, dilution, and reaction) by the same factor. As a  
 192 result, the attenuation factor  $\lambda$  scales the entire net-flux expression rather than only the  
 193 accumulation term. Because  $\lambda$  depends on surface condition, extended field use or residue  
 194 accumulation may alter its value, and periodic re-characterization is recommended to maintain  
 195 accurate flux estimates.

196 **Table S7.** The cumulative integrated differences between theoretical and measured values from  
 197 Figure S11 for both the filling and emptying phases, and the resulting  $\lambda$  values. While the signed  
 198 fill and empty values appear approximately balanced, their directional impact differs, and the  
 199 underlying asymmetry becomes critical in real measurements.

Gas	$\int (\text{Theo} - \text{Meas})_{\text{fill}}$	$\int (\text{Theo} - \text{Meas})_{\text{empty}}$	difference	$\lambda_{\text{fill}}$	$\lambda_{\text{empty}}$
$\text{NH}_3$	16.23	-16.33	-0.20	5.40	0.36
$\text{NO}_2$	13.40	-13.64	-0.23	3.04	0.40
HONO	13.57	-13.65	-0.05	3.14	0.40

200  
 201  
 202

## S8. References

- Burkholder, J. B., Sander, S. P., Abbatt, J. P. D., Barker, J. R., Cappa, C., Crouse, J. D., Dibble, T. S., Huie, R. E., Kolb, C. E., Kurylo, M. J., Orkin, V. L., Percival, C. J., Wilmout, D. M., Wine, P. H. (2020). Chemical Kinetics and Photochemical Data for Use in Atmospheric Studies - Evaluation Number 19. NASA JPL Publication 19-5. Tables 4C-1-2, 4C-1-3, and 4C-7-2. <https://jpldataeval.jpl.nasa.gov/>
- Crilley, L. R., Lao, M., Salehpoor, L., & VandenBoer, T. C. (2023). Emerging investigator series: an instrument to measure and speciate the total reactive nitrogen budget indoors: description and field measurements. *Environmental Science: Processes and Impacts*, 25(3), 389–404. <https://doi.org/10.1039/d2em00446a>
- Ebnesajjad, S. (2017). Introduction to Fluoropolymers. *Applied Plastics Engineering Handbook: Processing, Materials, and Applications: Second Edition*, 55–71. <https://doi.org/10.1016/B978-0-323-39040-8.00003-1>
- Ellis, R. A., Murphy, J. G., Pattey, E., Van Haarlem, R., O'brien, J. M., & Herndon, S. C. (2010). Characterizing a Quantum Cascade Tunable Infrared Laser Differential Absorption Spectrometer (QC-TILDAS) for measurements of atmospheric ammonia. *Atmos. Meas. Tech*, 3, 397–406.
- George, C., Ammann, M., D'Anna, B., Donaldson, D. J., & Nizkorodov, S. A. (2015). Heterogeneous Photochemistry in the Atmosphere. *Chemical Reviews*, 115(10), 4218–4258. <https://doi.org/10.1021/cr500648z>
- Lao, M., Crilley, L. R., Salehpoor, L., Furlani, T. C., Bourgeois, I., Andrew Neuman, J., Rollins, A. W., Veres, P. R., Washenfelder, R. A., Womack, C. C., Young, C. J., & VandenBoer, T. C. (2020). A portable, robust, stable, and tunable calibration source for gas-phase nitrous acid (HONO). *Atmospheric Measurement Techniques*, 13(11), 5873–5890. <https://doi.org/10.5194/amt-13-5873-2020>
- Ma, L., Pascalidou, E. M., Wiame, F., Zanna, S., Maurice, V., & Marcus, P. (2020). Passivation mechanisms and pre-oxidation effects on model surfaces of FeCrNi austenitic stainless steel. *Corrosion Science*, 167, 108483. <https://doi.org/10.1016/J.CORSCI.2020.108483>
- Pape, L., Ammann, C., Nyfeler-Brunner, A., Spirig, C., Hens, K., & Meixner, F. X. (2009). An automated dynamic chamber system for surface exchange measurement of non-reactive and reactive trace gases of grassland ecosystems. *Biogeosciences*, 6(3), 405–429. <https://doi.org/10.5194/bg-6-405-2009>
- Plake, D., Stella, P., Moravek, A., Mayer, J. C., Ammann, C., Held, A., & Trebs, I. (2015). Comparison of ozone deposition measured with the dynamic chamber and the eddy covariance method. *Agricultural and Forest Meteorology*, 206, 97–112. <https://doi.org/10.1016/j.agrformet.2015.02.014>

Salehpoor, L., & VandenBoer, T. C. (2023). Suppressor and calibration standard limitations in cation chromatography of ammonium and 10 alkylamines in atmospheric samples. *Analytical Methods*, *15*(31), 3822–3842. <https://doi.org/10.1039/d3ay01158e>

Sapag, K., Vallone, A., Blanco, A. A. G., & Solar, C. (2010). Adsorption of Methane in Porous Materials as the Basis for the Storage of Natural Gas. *Natural Gas*. <https://doi.org/10.5772/9846>