



1 **Storm-Time Asymmetries at Magnetic Conjugate Points: A Distribution-Aware**
2 **Benchmark for GNSS**

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7 **Abstract**

8 Geomagnetic storms disrupt the Global Navigation Satellite System (GNSS) and
9 transionospheric links through rapid asymmetric ionospheric variability. In this study, three
10 widely used empirical models (IRI-2016, IRI-Plas, and NeQuick2) were used against GNSS-
11 derived Total Electron Content (TEC) at two magnetic conjugate pairs (mid- and low-latitude)
12 during the geomagnetic storm of August 25–27, 2018. Rather than assessing storm-time
13 predictability, these models were employed as quiet-time reference baselines to quantify storm-
14 time deviations and hemispheric asymmetry. Model performance was evaluated using the Mean
15 Absolute Error (MAE), Root Mean Square Error (RMSE), and distribution-aware Kullback–
16 Leibler divergence (KLD). This study introduces a novel conjugate-point validation framework
17 augmented by KLD that uniquely captures both magnitude errors and structural distributional
18 mismatches between hemispheres. This is a critical aspect of GNSS reliability that is
19 overlooked by conventional metrics. The results indicate a phase-dependent performance: all
20 models exhibit degradation during the main phase, with the largest errors and structural
21 mismatches occurring at the equator. KLD reveals distributional distortions (variance,
22 skewness, tails) that MAE and RMSE cannot, particularly at the storm onset. NeQuick2
23 demonstrates superior performance only during the recovery phase, which is consistent with its
24 solar-flux-driven parameterization but limited topside representation. By integrating a
25 conjugate-point framework with distribution-aware validation, this study elucidates where
26 empirical baselines fail under storm conditions, and why hemispheric responses diverge. This
27 approach clarifies the model limitations relevant to GNSS reliability and motivates the
28 development of hybrid data-assimilative schemes that incorporate dynamic drivers while being
29 evaluated with both magnitude- and structure-sensitive metrics.

30 **Keywords.** Total Electron Content (TEC); GNSS; IRI-2016; NeQuick2; Conjugate Points;
31 Kullback-Leibler Divergence (KLD)



32 **Short Summary**

33 Ionospheric storms affect GNSS differently in conjugate hemispheres. Using GNSS-derived
34 TEC during the 25–27 Aug 2018 storm, we assess IRI-2016, IRI-Plas, and NeQuick2 as quiet-
35 time baselines. Magnitude errors (MAE, RMSE) and structure-sensitive KLD show phase-
36 dependent degradation, with equatorial asymmetries strongest. This framework reveals hidden
37 mismatches overlooked by conventional metrics and improves GNSS reliability assessment.

38 **1 Introduction**

39 The ionospheric response to geomagnetic storms remains a significant challenge in
40 space weather and wireless communication research because of its nonlinear dynamics
41 (Bojilova et al., 2024) and substantial impact on Global Navigation Satellite System (GNSS)
42 positioning, including scintillation-induced degradation (Aguiar et al., 2025; Yang et al., 2020),
43 satellite links, and transionospheric signal propagation (Li et al., 2024). It is important to note
44 that the objective of this study is not to evaluate the storm-time predictability of empirical
45 models; rather, these models are employed as quiet-time baselines to quantify storm-time
46 deviations and interhemispheric differences using distribution-aware metrics (Chou et al., 2023;
47 Peng et al., 2024). During geomagnetic disturbances, interactions between the solar wind and
48 magnetosphere inject energy into the ionosphere–thermosphere system through mechanisms
49 such as particle precipitation, prompt penetration of electric fields (PPEFs) (Tsurutani et al.,
50 2008), and disturbance dynamo currents (Lee et al., 2025; Paul et al., 2025; Venugopal et al.,
51 2025). These drivers disrupt the electrodynamic balance, alter the neutral composition, and
52 modulate global circulation, resulting in significant Total Electron Content (TEC, hereafter
53 VTEC when mapped to the vertical) variability that directly translates into signal delays, phase
54 scintillation, and GNSS service degradation (Aguiar et al., 2025; Bojilova et al., 2024; Li et al.,
55 2024; Yang et al., 2020).

56 A particularly incisive yet underutilized framework for investigating storm-time
57 dynamics is the analysis of magnetic conjugate points (Förster and Cnossen, 2013; Habarulema
58 et al., 2020; Laundal et al., 2017). These locations, connected by common geomagnetic field
59 lines, offer a natural laboratory for assessing the interhemispheric coupling (Habarulema et al.,
60 2020; Wang et al., 2021). Theoretically, such coupling should result in mirrored ionospheric
61 responses at the conjugate sites. However, empirical studies, including those by Dmitriev et al.
62 (Dmitriev et al., 2017), Habarulema et al. (Habarulema et al., 2020), and Wang et al. (Wang et
63 al., 2021), have demonstrated systematic non-mirrored behavior driven by hemispheric
64 differences in geomagnetic field geometry, dip angle, neutral composition, and thermospheric



winds (Habarulema et al., 2020; Laundal et al., 2017; Wang et al., 2021). Such non-mirrored behaviors are particularly critical for GNSS applications because prediction algorithms often assume hemispheric symmetry. These divergences can exceed approximately 100–150% during intense storms, such as the 2015 St. Patrick’s Day event (Astafyeva et al., 2015) and surpassed 200% in the American sector during the August 25–27, 2018 storm, as analyzed in subsequent multi-instrumental and modeling studies (Cherniak and Zakharenkova, 2022; Zhai et al., 2023), resulting in severe and unpredictable GNSS errors. While extreme asymmetries (>200%) have been reported in the American sector during this storm (Cherniak and Zakharenkova, 2022; Zhai et al., 2023), our study focuses on the European, African, and Asian sectors to provide a complementary global perspective and assess whether such severe discrepancies are also prevalent in other longitudinally distinct regions with high user density. Despite their practical importance, conjugate-point discrepancies remain insufficiently quantified in terms of both their physical drivers and their representation in operational models (Cnossen and Förster, 2016; Laundal et al., 2017). Global model validations (e.g., Bilitza et al., 2017 (Bilitza et al., 2017)) often emphasize averaged metrics, whereas regional case studies, such as Sherif et al. (2024) (Sherif et al., 2024), highlight that performance can vary substantially across local longitudinal sectors, underscoring the need for conjugate-point frameworks that explicitly capture hemispheric differences.

Despite the recognized need for improvement, empirical models such as IRI-2016, IRI-Plas, and NeQuick2 continue to serve as standard reference frameworks for ionospheric specifications in GNSS and communication system applications (Chou et al., 2023; Luo et al., 2023; Peng et al., 2024). These models are particularly effective in replicating climatological behavior under quiet time conditions. However, they were not originally developed to account for storm-time electrodynamics (Cnossen and Förster, 2016; Laundal et al., 2017). This limitation has been acknowledged for some time, yet validation studies predominantly focus on global averages or pointwise error statistics, such as RMSE or MAE (Bilitza et al., 2017; Luo et al., 2023; Sherif et al., 2024). Although these metrics effectively quantify magnitude errors, they do not adequately capture distributional differences in TEC variability, structural mismatches across hemispheres, or the storm-phase dependence of model skill (Jiang et al., 2024). Moreover, very few validation efforts have explicitly employed a conjugate-point framework (Astafyeva et al., 2015; Habarulema et al., 2020; Laundal et al., 2017; Wang et al., 2021; Zhai et al., 2023), leaving hemispheric divergence, a critical aspect of storm-time GNSS degradation, largely unexplored.



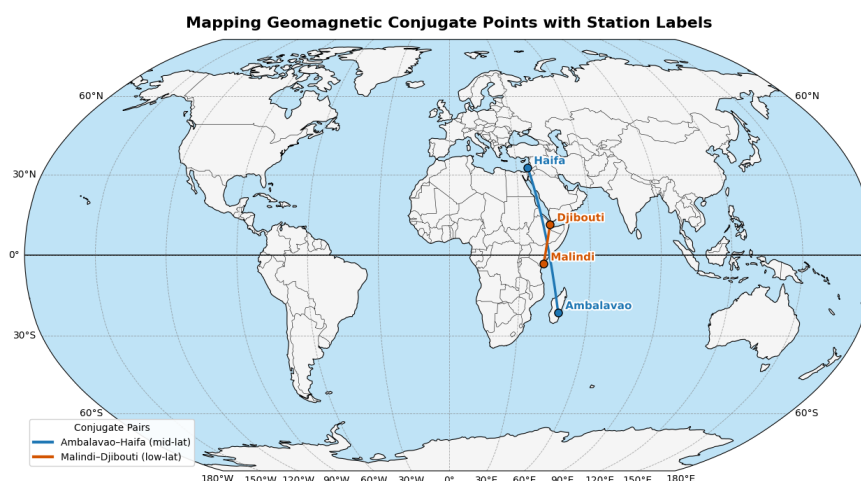
98 Utilizing GNSS-derived TEC as a widely recognized benchmark, we employed three
99 complementary validation metrics: (i) Mean Absolute Error (MAE), (ii) root-mean-square error
100 (RMSE), and (iii) distribution-sensitive Kullback–Leibler Divergence (KLD). While MAE and
101 RMSE assess the pointwise accuracy (Bilitza et al., 2017; Tang et al., 2020), KLD evaluates
102 the structural discrepancies between the modeled and observed TEC distributions (Adolfs et
103 al., 2022). In this study, KLD was calculated between the probability distributions of the
104 modeled and observed TEC during the storm phases (initial, main, and recovery) using
105 consistent binning across both datasets (refer to Section 2.4 for implementation details). This
106 metric was specifically selected because of its asymmetry and sensitivity to differences in the
107 tails of distributions, which are essential for identifying storm-time anomalies that manifest as
108 rare but high-impact events in the TEC. Given that the TEC structure directly influences range
109 delay and scintillation risk, distribution-aware validation is operationally pertinent to GNSS
110 integrity (Bojilova et al., 2024; Li et al., 2024; Martinon et al., 2023). Beyond geospace science,
111 similar distribution-sensitive divergence metrics are extensively utilized in wireless
112 communication research to evaluate channel reliability and error propagation (Al-Jarrah et al.,
113 2023; Manimegalai and Bhagyaveni, 2019), highlighting their methodological significance in
114 detecting storm-time TEC anomalies. This cross-disciplinary application underscores the
115 methodological robustness of incorporating KLD into our framework, as both ionospheric TEC
116 variability and wireless channel fluctuations exhibit similar distribution-sensitive error
117 characteristics.

118 By integrating these approaches, our analysis (i) elucidates the conditions under which
119 empirical models fail during storm forcing, (ii) quantifies cross-hemispheric differences at
120 conjugate locations, and (iii) underscores the methodological benefits of distribution-aware
121 validation for communication-focused applications (Adolfs et al., 2022; Al-Jarrah et al., 2023;
122 Habarulema et al., 2020; Jiang et al., 2024; Luo et al., 2023; Manimegalai and Bhagyaveni,
123 2019; Wang et al., 2021). In doing so, we reconceptualize the role of empirical ionospheric
124 models in storm-time contexts, not as predictive tools for disturbed conditions, but as reference
125 baselines whose limitations, when properly characterized, provide critical insights into GNSS
126 reliability and wireless system vulnerability. This dual perspective, combining conjugate-point
127 analysis with distribution-sensitive metrics, offers a novel contribution by bridging space
128 weather science and applied wireless communication (Adolfs et al., 2022; Al-Jarrah et al., 2023;
129 Habarulema et al., 2020; Jiang et al., 2024; Luo et al., 2023; Manimegalai and Bhagyaveni,
130 2019; Wang et al., 2021).



131 2 Data and Model

132 This study examines the ionospheric response during a geomagnetic storm from August
133 25 to 27, 2018, at two pairs of magnetic conjugate stations: Ambalavao (21.9°S, 46.8°E) and
134 Haifa (32.8°N, 35.0°E), representing mid-latitudes, and Malindi (3.0°S, 40.1°E) and Djibouti
135 (11.5°N, 42.9°E), representing low latitudes (Akala et al., 2021; Astafyeva et al., 2022; Piersanti
136 et al., 2020; Zhai et al., 2023). These stations were strategically selected to capture distinct
137 geophysical regimes, facilitating a systematic evaluation of hemispheric asymmetries across
138 both mid- and equatorial latitudes (Förster and Haaland, 2015; Laundal and Richmond, 2017).
139 Figure 1 illustrates the geographic distribution of stations and their conjugate linkages. The East
140 African and Eastern Mediterranean sectors were chosen because of their high density of GNSS
141 users and critical satellite communication routes, making the precise characterization of TEC
142 variability in these regions essential for service reliability (European GNSS Service Centre
143 (GSC), 2015; Ionospheric Correction Algorithms (NeQuick-G / NTCM-G), 2025). This
144 configuration is particularly pertinent for GNSS users because spatial and hemispheric
145 variability in the TEC directly results in positioning errors, signal delays, and service
146 degradation in wireless communication systems.



147

148 **Fig. 1** Geographic distribution of GNSS stations and their corresponding magnetic conjugate
149 links. Station identifiers (IGS codes) are indicated, and great-circle arcs are utilized to
150 approximate the connectivity of the geomagnetic field lines. The latitude and longitude
151 graticules were presented at intervals of 20°.

152 2.1 Observational Data



153 Total Electron Content (TEC) was derived from dual-frequency GNSS data obtained via the
154 International GNSS Service (IGS). RINEX-format observations from GPS, GLONASS, and
155 Galileo satellites (when available) were processed at a 30-second resolution using the
156 IONOLAB-TEC software suite (Sezen et al., 2013). For each epoch, all satellite–receiver links
157 above a 30° elevation mask were included. The vertical TEC (VTEC) values were computed
158 per satellite at the ionospheric pierce point and subsequently averaged to obtain a station-level
159 time series, thereby reducing satellite-specific noise while retaining the storm-time variability
160 most relevant to the GNSS navigation accuracy.

161 To ensure calibration accuracy, the IONOLAB-BIAS algorithm was employed to
162 estimate and eliminate differential code biases (DCBs) from both receivers and satellites
163 (Arikan et al., 2008). This process incorporates precise IGS orbits and IONEX-derived global
164 DCB products (Hernández-Pajares et al., 2009). The methodology has undergone extensive
165 validation against global IGS centers, and is recognized for providing reliable single-station
166 TEC retrievals under both quiet and disturbed conditions. Although GNSS-derived TEC is
167 widely regarded as the reference ground truth, it is important to acknowledge the limitations
168 associated with receiver biases, multipaths, and calibration errors. These effects were mitigated
169 through IONOLAB-BIAS correction and validated orbits/DCB products, ensuring that the
170 residual errors were minimal relative to the storm-time TEC variability of interest in this study,
171 as supported by the comprehensive validation of IONOLAB methodology against global IGS
172 products (Sezen et al., 2013). Station metadata, receiver types, and data availability windows
173 are summarized in Table S1 (Supplementary Information).

174 **2.2 Empirical Models**

175 Three widely used empirical ionospheric models were selected for this study: IRI-2016,
176 IRI-Plas, and NeQuick2 (Bilitza et al., 2017; Gulyaeva et al., 2011; Nava et al., 2008). These
177 models are not intended as predictive tools for storm-time conditions, as they were not designed
178 for such purposes. Instead, they serve as climatological baselines, against which deviations
179 during storm-time and hemispheric asymmetries can be systematically quantified. From the
180 perspective of wireless communication, these models are frequently integrated into GNSS error
181 correction, integrity monitoring, and service reliability assessments, rendering their evaluation
182 particularly pertinent under disturbed conditions (European GNSS Service Centre (GSC), 2015;
183 Ionospheric Correction Algorithms (NeQuick-G / NTCM-G), 2025).

184 IRI-2016 represents the climatological monthly medians of ionospheric parameters
185 (Bilitza et al., 2017). IRI-Plas extends IRI by incorporating plasmaspheric contributions to the



186 Total Electron Content (TEC), thereby enhancing the representation of the topside ionosphere
187 (Gulyaeva et al., 2011). NeQuick2 is a semi-empirical model driven by solar flux inputs (F10.7,
188 or the effective ionization level, A_z) to adjust the electron density profiles dynamically (Nava
189 et al., 2008). Although storm-time empirical models exist (e.g., the IRI-based STORM
190 extension) (Araujo-Pradere et al., 2002), this study specifically focuses on the performance of
191 the standard, unmodified versions of IRI-2016, IRI-Plas, and NeQuick2. These standard
192 versions are the most widely implemented in operational GNSS receivers and communication
193 systems for real-time error correction. Therefore, evaluating their inherent limitations during
194 storms is of direct practical relevance (European GNSS Service Centre (GSC), 2015;
195 Ionospheric Correction Algorithms (NeQuick-G / NTCM-G), 2025).

196 2.3 Storm Phases

197 Model outputs and GNSS observations were systematically compared across three
198 distinct storm phases, initial, main, and recovery, as defined by the temporal variation in the
199 Dst index (Gonzalez et al., 1994). This phase-resolved methodology facilitates the assessment
200 of model performance in relation to storm progression, rather than averaging results across
201 dynamically disparate intervals. This distinction is crucial for relevance to communication
202 systems, as GNSS errors typically reach their peak during the main phases and stabilize during
203 the recovery phase.

204 2.4 Statistical Metrics

205 Three statistical measures were employed to assess the model performance relative to
206 GNSS-derived TEC:

$$207 \quad MAE = \frac{1}{n} \sum_{i=1}^n |TEC_i - TEC_p| \quad (1)$$

$$208 \quad RMSE = \sqrt{\frac{1}{n} \sum_{i=1}^n (TEC_i - TEC_p)^2} \quad (2)$$

209 where TEC_i is the observed value, TEC_p the model output, and n the number of samples.
210 Equations (1) and (2) capture pointwise errors, with MAE emphasizing the average deviations
211 and RMSE amplifying larger mismatches.

212 To go beyond magnitude-based errors, we introduced the Kullback–Leibler Divergence (KLD):

$$213 \quad KLD = \sum_i P(i) \log \frac{P(i)}{Q(i)} \quad (3)$$



214 In this context, $P(i)$ and $Q(i)$ denote the probability distributions of the observed and modeled
215 Total Electron Content (TEC), respectively, with the divergence of the model distribution Q
216 from the observed distribution P being the focus. A lower Kullback-Leibler divergence (KLD)
217 signifies closer alignment in the distributional structure. Unlike Equations (1) and (2), Equation
218 (3) is sensitive to discrepancies in variance, skewness, and multimodality, which often become
219 pronounced during storm-time dynamics (Kullback and Leibler, 1951; Sason and Verdu, 2016).
220 For implementation, probability distributions were estimated using histograms with bin widths
221 determined by the Freedman–Diaconis rule to ensure consistency across datasets (Freedman
222 and Diaconis, 1981). Kernel density estimation (KDE) was also evaluated as part of a
223 comprehensive sensitivity analysis. To prevent undefined values in the logarithmic term, a
224 small smoothing constant ($\epsilon \approx 10^{-6}$) was added to all bins following standard divergence-
225 analysis practice (Cover and Thomas, 2005). This setup is also consistent with wireless
226 communication research, where divergence metrics are commonly applied to evaluate the
227 channel reliability and error propagation.

228 A comprehensive sensitivity analysis was performed to evaluate the robustness of the
229 Kullback-Leibler divergence (KLD) metric in relation to the inherent subjectivity of probability
230 density estimation. KLD was recalculated for each station-model pair across a broad spectrum
231 of histogram bin counts (ranging from 20 to 60) and Gaussian kernel density estimation (KDE)
232 bandwidths (spanning 0.30 to 0.70 TECU). Variability in the results was quantified using the
233 relative percentage change ($\Delta\%$) across these parameter ranges. This analysis demonstrated that
234 KLD is systematically sensitive to these parameters; it generally increases with finer histogram
235 binning and decreases with broader KDE smoothing. Importantly, the extent of this sensitivity
236 is highly contingent on both the model and the geophysical location.

237 Across all stations, the Δ bin values ranged from 3.68% to 75.64%, whereas the Δ_{KDE}
238 values varied from 10.85% to 67.23%. The mid-latitude stations, Ambalavao and Djibouti,
239 exhibited significant sensitivity, with the KLD increasing consistently with finer binning
240 ($\Delta_{\text{bins}} \approx 61\text{--}67\%$ across models) and decreasing with broader KDE smoothing ($\Delta_{\text{KDE}} \approx 56\text{--}$
241 64% at Ambalavao; $11\text{--}58\%$ at Djibouti). Haifa demonstrated the highest parameter sensitivity,
242 particularly for NeQuick2 ($\Delta_{\text{bins}} \approx 75.6\%$, $\Delta_{\text{KDE}} \approx 67.2\%$), whereas IRI-PLAS remained
243 relatively stable at this location ($\Delta_{\text{bins}} \approx 35.8\%$, $\Delta_{\text{KDE}} \approx 30.0\%$). Conversely, the equatorial
244 station Malindi was an outlier; its long-tailed PDFs rendered the KLD nearly insensitive to
245 binning ($\Delta_{\text{bins}} \approx 3\text{--}7\%$ across models), although it remained responsive to the KDE
246 bandwidth ($\Delta_{\text{KDE}} \approx 45\text{--}48\%$ for IRI-PLAS and NeQuick2).



247 Crucially, the divergence between the model and observations remained consistent
248 across all parameter configurations. Consequently, our conclusions based on KLD are
249 indicative of authentic structural differences rather than artifacts arising from histogram or KDE
250 settings. The comprehensive results of this sensitivity analysis are presented in Supplementary
251 Figures S1-S12 and summarized in Supplementary Table S2. This thorough validation
252 emphasizes the robustness of our distribution-aware evaluation and offers an essential context
253 for interpreting the KLD results discussed in the subsequent sections.

254 **2.5 Theoretical Framework**

255 Conjugate-point analysis is based on storm-time electrodynamics. Prompt penetration
256 electric fields (PPEFs) and disturbance dynamo currents facilitate large-scale plasma
257 redistribution along the magnetic field lines (Blanc and Richmond, 1980; Tsurutani et al.,
258 2008). Although theoretical models anticipate mirrored behaviors at conjugate points, empirical
259 observations consistently reveal hemispheric asymmetries owing to geomagnetic field
260 geometry, conductivity imbalances, and neutral wind filtering (Förster and Haaland, 2015;
261 Laundal and Richmond, 2017). Total Electron Content (TEC) serves as a crucial diagnostic tool
262 for these processes, and its distribution across hemispheres provides a rigorous test for empirical
263 models. Beyond its geophysical importance, this framework elucidates the mechanisms by
264 which ionospheric storms impair Global Navigation Satellite System (GNSS) positioning
265 accuracy and communication reliability, underscoring the necessity of cross-disciplinary
266 evaluation methods. The theoretical expectation of both coupling and asymmetry directly
267 informs our methodological choices: selecting conjugate pairs offers a controlled experimental
268 setup, whereas the Kullback-Leibler divergence (KLD) metric is particularly suited for
269 detecting the non-mirrored distributional differences predicted by theory (Bilitza et al., 2017).
270 Consequently, our framework provides a direct observational test for these electrodynamic
271 principles. Therefore, precise quantification of hemispheric divergence, a critical factor in
272 storm-time GNSS degradation, remains a central challenge, emphasizing the need for
273 conjugate-point validation frameworks. As conjugate sites do not share a common local time
274 (LT) at a given UT, the interpretation of hemispheric asymmetry is potentially confounded by
275 diurnal illumination. In future analyses, we will employ LT-aligned comparisons and explicit
276 LT controls to isolate electrodynamic coupling from diurnal effects.

277 **3 Results**

278 **3.1 Storm Overview and Ionospheric Context**

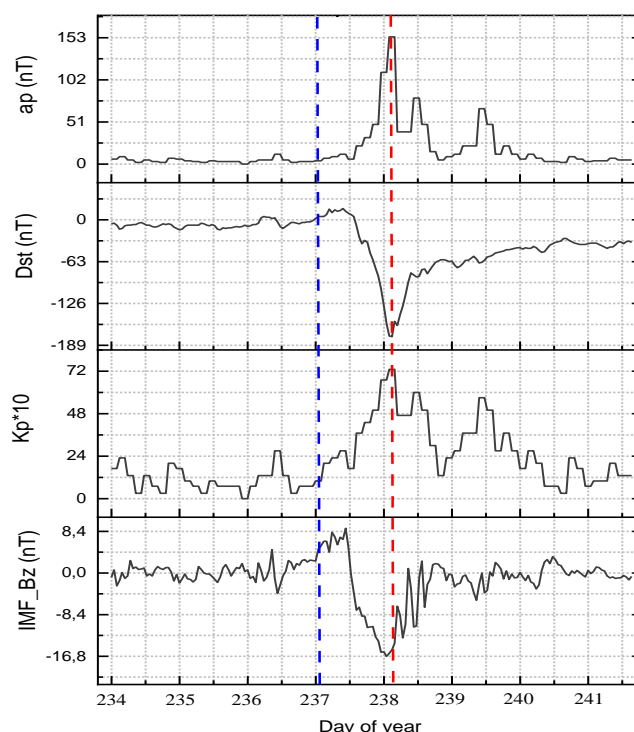


279 The geomagnetic storm examined in this study occurred between August 25 and 27,
280 2018 and exhibited characteristics of a moderately intense to strong event. As illustrated in
281 Figure 2, Dst reached -176 nT, Kp peaked at 7.3, and a_p increased to 154 nT (Piersanti et al.,
282 2020). The main phase, characterized by a prolonged southward turning of the IMF Bz,
283 extended from 18:00 UT on August 25 to 06:00 UT on August 26 (Piersanti et al., 2020). Taken
284 together, these indicators are consistent with strong storm conditions (Piersanti et al., 2020).

285 Storm classifications vary in the literature. Akala et al. (2021) documented that the
286 geomagnetic activity on August 26, 2018, was consistent with a G3-level event precipitated by
287 a solar filament eruption on August 20 (Akala et al., 2021). Subsequent analyses suggested that
288 the storm was predominantly influenced by weak CME transients interacting with Corotating
289 Interaction Regions (CIRs) and High-Speed Streams (HSSs) rather than by a strong magnetic
290 cloud, which is typically associated with severe geomagnetic disturbances (Akala et al., 2021;
291 Gonzalez et al., 2011).

292 Although Prompt Penetration Electric Fields (PPEFs) were relatively modest (≈ 0.3
293 mV/m; Akala et al., 2021) (Akala et al., 2021), our observations reveal significant ionospheric
294 responses, particularly in the form of pronounced Total Electron Content (TEC) deviations at
295 magnetic conjugate points. This underscores two critical aspects: (i) even moderately driven
296 storms with sustained southward Interplanetary Magnetic Field (IMF) Bz can induce strong
297 interhemispheric ionospheric coupling and (ii) localized spatial variability plays a decisive role
298 in shaping storm-time ionospheric impacts.

299 Consequently, while the August 2018 event may not be classified as an extreme storm
300 based on solar drivers, its impact on the ionosphere serves as a valuable case study for assessing
301 the performance of the empirical model under disturbed and complex geomagnetic conditions.
302 This perspective is consistent with previous research indicating that even G2–G3 storms can
303 significantly disrupt GNSS positioning and communication systems when they occur under
304 favorable local time and background thermospheric conditions (Valdés-Abreu et al., 2021; Xue
305 et al., 2024).



306

307 **Fig. 2** Temporal variations in the IMF Bz, Dst, Kp, and ap indices observed from August 22 to
308 29, 2018. The main phase of the storm, which occurred from August 25, 18:00 UT to August
309 26, 06:00 UT, is delineated by the vertical dashed lines. A pronounced minimum in the Dst
310 index and sustained southward orientation of the IMF Bz corroborated the intensity of the
311 storm.

312 3.2 TEC Model Evaluation at Conjugate Points

313 To assess the efficacy of empirical ionospheric models in representing storm-time
314 dynamics, we conducted a comparative analysis of NeQuick2, IRI-2016, and IRI-Plas during a
315 geomagnetic storm that occurred from August 25 to 27, 2018. This evaluation concentrated on
316 two magnetic conjugate station pairs, Ambalavao–Haifa and Malindi–Djibouti, encompassing
317 both mid-latitude and equatorial responses. Model outputs were systematically compared with
318 GNSS-derived Total Electron Content (TEC) using three complementary metrics: MAE,
319 RMSE, and KLD. This hybrid framework, which integrates magnitude-based errors with
320 distribution-sensitive divergence, offers a comprehensive assessment of the model performance
321 across various storm phases. This approach is particularly valuable given the nonlinear,
322 asymmetric, and rapidly evolving nature of ionospheric disturbances during geomagnetic



storms, particularly when interhemispheric processes are considered (Bilitza et al., 2017; Cherniak and Zakharenkova, 2019; Vankadara et al., 2022).

Although these empirical models are extensively employed for climatological specifications, they possess inherent limitations in accurately representing the rapid electrodynamic processes that occur during geomagnetic disturbances. These processes include the prompt penetration of electric fields, storm-time thermospheric wind surges, and hemispheric conductivity gradients (Blanc and Richmond, 1980; Tsurutani et al., 2008).

3.2.1 Ambalavao–Haifa pair: How do the models behave at mid-latitudes?

As illustrated in Table 1, all three models demonstrated optimal performance during the recovery phase, with NeQuick2 exhibiting the lowest errors (MAE \approx 0.4 TECU; RMSE \approx 1.0 TECU). Conversely, the error rates increased significantly during the main phase, particularly for IRI-2016 (while IRI-Plas occasionally yielded a smaller MAE at mid-latitudes during the MP; Table 1), highlighting their limited capacity to accurately capture rapid storm-time fluctuations. These findings align with previous reports indicating that IRI models tend to underestimate the daytime TEC and oversimplify storm variability (Bilitza et al., 2017; Cherniak and Zakharenkova, 2019; Endeshaw, 2020; Gulyaeva et al., 2011).

Table 1 TEC model performance for Ambalavao–Haifa across storm phases (BP, MP, and RP), errors (MAE, RMSE, and TECU), and divergence (KLD; –) are reported relative to OBS $|\Delta\text{TEC}|$.

Metric	Phase	Observed (Amb–Haifa)	NeQuick2	IRI-2016	IRI-Plas
MAE (TECU)	BP	6.0	1.7	1.8	2.1
	MP	6.2	2.8	2.3	1.3
	RP	2.3	0.4	0.5	1.5
RMSE (TECU)	BP	6.4	1.9	2.1	2.8
	MP	6.4	3.0	2.5	1.7
	RP	4.2	1.0	1.3	3.5
KLD (–)	BP	0.08	0.30	0.31	0.17
	MP	0.07	0.22	0.26	0.13
	RP	0.16	0.26	0.31	0.16

Note: MAE and RMSE are in TECU, and KLD is dimensionless.

An important insight derived from the Kullback–Leibler Divergence (KLD) is that elevated values during the initial phase indicate that despite a moderate Mean Absolute Error (MAE) and Root Mean Square Error (RMSE), the models inaccurately represented the statistical structure of the Total Electron Content (TEC) distribution. Variance, skewness, and tail behavior, which are often obscured by pointwise metrics, have proven crucial in capturing the onset of storm-driven changes (Arikan et al., 2016; Karatay et al., 2010). To our knowledge, this is among the earliest distribution-aware (KLD) validations explicitly framed within a conjugate-point analysis for this storm and station set.



3.2.2 Haifa–Ambalavao Reverse Evaluation: The Hemispheric Disconnect

Upon reversing the evaluation to compare Haifa with Ambalavao, distinct hemispheric asymmetry was evident (Table 2). Errors were consistently more pronounced at the Northern Hemisphere site, particularly during the main phase, suggesting increased modeling challenges at higher northern mid-latitudes. This phenomenon aligns with variations in the seasonal solar illumination, geomagnetic field geometry, and background electron density gradients (Laundal et al., 2017; Mendillo, 2006).

Table 2 Directional (reverse) evaluation for Haifa–Ambalavao across BP, MP, RP; errors (MAE, RMSE; TECU) and divergence (KLD; –) reported relative to OBS $|\Delta\text{TEC}|$.

Metric	Phase	Observed (Haifa–Amb)	NeQuick2	IRI-2016	IRI-Plas
MAE (TECU)	BP	6.0	2.8	4.3	2.3
	MP	6.2	4.6	6.6	2.6
	RP	2.3	0.5	1.2	0.7
RMSE (TECU)	BP	6.4	3.3	4.8	2.8
	MP	6.4	4.7	6.7	2.9
	RP	4.2	1.2	2.2	1.4
KLD (–)	BP	0.08	0.04	0.05	0.03
	MP	0.07	0.08	0.07	0.06
	RP	0.16	0.11	0.07	0.10

Note: MAE and RMSE are in TECU, and KLD is dimensionless.

The Kullback-Leibler divergence (KLD) further accentuates the asymmetry, indicating that the models encountered difficulties with the statistical structure of Total Electron Content (TEC) distributions in Haifa, beyond merely the mean errors.

3.2.3 Malindi–Djibouti Pair: The Challenge of the Equator

For the Malindi–Djibouti pair, performance declined further (Table 3). During the main phase, the IRI-2016 model exhibited a Mean Absolute Error (MAE) exceeding 15 Total Electron Content Units (TECU), with a Root Mean Square Error (RMSE) approaching 16 TECU, which is significantly larger than the errors observed at mid-latitudes. These discrepancies underscore the inherent complexity of the equatorial ionosphere, which is significantly influenced by the Equatorial Ionization Anomaly (EIA), neutral wind reversal, and storm-time coupling processes (Cherniak and Zakharenkova, 2022; Zhang et al., 2021).

Table 3 TEC model performance for Djibouti–Malindi across BP, MP, and RP; errors (MAE, RMSE, TECU) and divergence (KLD; –) reported relative to OBS $|\Delta\text{TEC}|$.

Metric	Phase	Observed (Mal–Dji)	NeQuick2	IRI-2016	IRI-Plas
MAE (TECU)	BP	12.5	10.5	10.9	5.5
	MP	13.3	15.7	15.2	8.1
	RP	5.3	5.0	4.8	2.4
RMSE (TECU)	BP	12.9	11.1	11.5	7.2
	MP	14.0	15.9	15.2	8.6



	RP	8.9	8.6	8.2	4.7
KLD (–)	BP	0.10	0.23	0.28	0.15
	MP	0.14	0.13	0.17	0.09
	RP	0.07	0.14	0.16	0.09

373 *Note: MAE and RMSE are in TECU, and KLD is dimensionless.*

374 The KLD values indicate that the model not only inaccurately estimated TEC
375 magnitudes, but also failed to maintain distributional characteristics, emphasizing the specific
376 challenges associated with equatorial dynamics.

377 3.2.4 NeQuick2's Edge in the Recovery Phase

378 Despite its recognized limitations, NeQuick2 consistently surpassed IRI-2016 and IRI-
379 Plas during the recovery phase, as shown in Table 4. With mean absolute error (MAE) values
380 below 1 total electron content unit (TECU), relatively low root mean square error (RMSE), and
381 Kullback-Leibler divergence (KLD), NeQuick2 demonstrated considerable robustness once the
382 ionosphere stabilized following the peak disturbance (Montenbruck and González Rodríguez,
383 2020; Nava et al., 2008).

384 **Table 4** Directional (reverse) evaluation for Malindi–Djibouti across BP, MP, and RP; errors
385 (MAE, RMSE, TECU) and divergence (KLD; –) reported relative to OBS $|\Delta\text{TEC}|$.

Metric	Phase	Observed (Mal–Dji)	NeQuick2	IRI-2016	IRI-Plas
MAE (TECU)	BP	12.5	3.9	4.2	10.7
	MP	13.3	2.9	3.3	8.7
	RP	5.28	0.7	1.2	3.8
RMSE (TECU)	BP	12.92	5.0	5.2	12.2
	MP	13.9	4.2	3.8	10.2
	RP	8.9	1.5	2.6	7.3
KLD (–)	BP	0.10	0.23	0.28	0.15
	MP	0.14	0.13	0.17	0.09
	RP	0.17	0.14	0.16	0.09

386 *Note: MAE and RMSE are in TECU, and KLD is dimensionless.*

387 The KLD values remained relatively low, indicating that NeQuick2 maintained its
388 structural integrity during the recovery. This relative success is likely attributable to its solar-
389 flux-driven parameterization, which facilitates dynamic adjustment during recovery, in contrast
390 to IRI-based models that rely on climatological medians and lack flexibility. However,
391 NeQuick2's simplified plasmaspheric representation resulted in an underestimation of the
392 topside TEC (Cherniak and Zakharenkova, 2019), highlighting that even the most reliable
393 empirical models have inherent limitations under storm conditions (Kashcheyev and Nava,
394 2019; Pezzopane et al., 2023; Pezzopane and Pignalberi, 2019).

395 4 Discussion

396 4.1 Model Limitations and Benchmark Role



397 In this study, empirical models were employed as baselines for quiet-time conditions to assess
398 deviations during storm periods; however, they were not used as predictors of storm events.
399 This analysis underscores a well-established limitation: empirical ionospheric models such as
400 IRI-2016, IRI-Plas, and NeQuick2 are not designed to predict storm-time phenomena (Bilitza
401 et al., 2017; Gulyaeva et al., 2011; Nava et al., 2008). These models are based on climatological
402 averages and are not intended to capture transient electrodynamic forcings (Bilitza et al., 2017;
403 Blanc and Richmond, 1980; Nava et al., 2008). Rather than viewing this as a deficiency, our
404 study reinterprets their function by providing benchmarks for quiet-time conditions against
405 which storm-time deviations can be systematically quantified. Extensions such as the IRI-based
406 STORM module aim to parameterize storm effects; however, they remain constrained by the
407 complexity of magnetosphere–ionosphere–thermosphere interactions (Araujo-Pradere et al.,
408 2002; Astafyeva et al., 2015; Blanc and Richmond, 1980). Therefore, evaluating baseline
409 empirical models under storm conditions is scientifically valuable not as a test of forecasting
410 ability but as a means to diagnose where physical processes diverge from climatological
411 assumptions (Bilitza et al., 2017; Gulyaeva et al., 2011). Notably, IRI-Plas occasionally
412 exhibited lower magnitude errors (MAE/RMSE) during the main phase at mid-latitudes (e.g.,
413 Tables 1 and 2), which may have resulted from its plasmaspheric extension stabilizing topside
414 contributions under certain conditions or from differences in parameterization. This
415 occasionally lower MAE/RMSE during the main phase is physically plausible: the
416 plasmaspheric extension can partially stabilize topside contributions and reduce magnitude
417 errors when the scale heights increase rapidly. However, this advantage does not consistently
418 translate into superior distributional fidelity, as indicated by the KLD, underscoring that IRI-
419 Plas may match totals while still misrepresenting variance, skewness, or tails. However,
420 establishing generality requires further case studies (Arikan et al., 2016; Gulyaeva et al., 2011).

421 **4.2 Hemispheric Asymmetry and Physical Insights**

422 One of the most evident outcomes is the pronounced hemispheric asymmetry observed
423 at the conjugate points, notably in the Malindi–Djibouti pair. Despite their magnetic linkage,
424 these sites exhibited significantly different TEC signatures during the main phase, with IRI-
425 2016 errors exceeding 15 TECU at Djibouti. This asymmetry cannot be attributed solely to
426 model deficiencies; it reflects genuine geophysical influences, including solar illumination
427 gradients, thermospheric wind circulation, and geographic offsets (Cnossen and Förster, 2016;
428 Förster and Cnossen, 2013; Förster and Haaland, 2015; Laundal et al., 2017; Laundal and
429 Richmond, 2017). These findings are consistent with previous reports on conjugate divergence



430 and underscore the necessity for conjugate frameworks in future validation studies because
431 assumptions of symmetry obscure significant physical variability (Cherniak and Zakharenkova,
432 2022; Habarulema et al., 2020; Zhang et al., 2021). Although the hemispheric asymmetries
433 discussed above reflect authentic geophysical drivers, conjugate sites do not share the same
434 local time (LT) at a given UT. Because TEC is strongly influenced by solar illumination (LT),
435 a portion of the observed divergence may result from LT offsets rather than solely from
436 hemispheric electrodynamics. Future analyses should incorporate LT-aligned composites or
437 explicit LT controls to isolate this contribution, which represents a limitation of this study
438 (Cnossen and Förster, 2016; Förster and Cnossen, 2013; Förster and Haaland, 2015; Laundal et
439 al., 2017; Laundal and Richmond, 2017; Mendillo, 2006).

440 **4.3 Structural Reliability via KLD**

441 Traditional error metrics, such as the Mean Absolute Error (MAE) and Root Mean Square Error
442 (RMSE), evaluate the mean accuracy but do not ascertain whether the statistical structure of
443 storm-time Total Electron Content (TEC) is preserved. By utilizing the Kullback–Leibler
444 Divergence (KLD), we identified instances where models appeared satisfactory in terms of
445 mean error yet inaccurately represented variance, skewness, or distribution tails, particularly at
446 the onset of storms (Arikan et al., 2016; Karatay et al., 2010; Kullback and Leibler, 1951). This
447 distinction is significant; a model that accurately predicts numerical values may still
448 inadequately capture the dynamics governing ionospheric variability. Our findings demonstrate
449 that KLD complements conventional metrics by revealing hidden discrepancies in model
450 performance (Arikan et al., 2016; Cover and Thomas, 2005; Karatay et al., 2010; Kullback and
451 Leibler, 1951; Sason and Verdu, 2016). The application of KLD to conjugate-point analysis
452 represents a methodological advancement beyond standard evaluation (Kullback and Leibler,
453 1951; Sason and Verdu, 2016). Operationally, an elevated KLD indicates that the model
454 underrepresents the distribution tails that drive rare but high-impact TEC excursions, which are
455 the most challenging for Global Navigation Satellite System (GNSS) correction/integrity and
456 wireless links. It is important to note that KLD quantifies the similarity of distributional shape
457 (variance, skewness, multimodality) rather than the absolute magnitude of errors; thus, a station
458 can exhibit large MAE/RMSE yet moderate KLD if the modeled distribution preserves shape,
459 whereas smaller magnitude errors can still yield higher KLD when the distributional structure
460 is incorrect—hence our joint reporting of MAE/RMSE and KLD.

461 Sensitivity analyses (Supplementary Figs. S1–S12; Tables S1–S2) confirm that these
462 conclusions are not artifacts of density estimation settings. The analysis revealed that NeQuick2



463 was the most parameter-sensitive model, as indicated by the significant fluctuations in the KLD
464 values depending on the bin count and KDE bandwidth. In contrast, IRI-Plas demonstrated
465 greater stability. This finding underscores that distribution-aware metrics not only identify
466 structural discrepancies but also highlight the vulnerability of different models to
467 methodological assumptions. Therefore, distribution-aware metrics are directly relevant to risk-
468 focused validation (Al-Jarrah et al., 2023; Kullback and Leibler, 1951; Manimegalai and
469 Bhagyaveni, 2019; Sason and Verdu, 2016).

470 **4.4 Future Directions**

471 The degradation observed during storm-time was most pronounced during the main
472 phase, wherein all models failed to accurately capture rapid electrodynamic forcing (Astafyeva
473 et al., 2015, 2022). Conversely, the recovery phase exhibited variability contingent upon the
474 model used, with NeQuick2 demonstrating a relative advantage (Cherniak and Zakharenkova,
475 2019; Kashcheyev and Nava, 2019; Montenbruck and González Rodríguez, 2020; Pezzopane
476 et al., 2023; Pezzopane and Pignalberi, 2019). This disparity underscores the necessity of
477 developing models capable of dynamic adaptation across different phases rather than relying
478 on static climatological assumptions (Adolfs et al., 2022; Jiang et al., 2024; Luo et al., 2023).

479 The most pronounced cross-hemispheric divergences are observed at equatorial
480 conjugate pairs, where complex factors such as pre-reversal enhancement (Fejer, 2002), EIA
481 modulation, and neutral wind asymmetries intensify the discrepancies between models and
482 observations (Cherniak and Zakharenkova, 2022; Cnossen and Förster, 2016; Habarulema et
483 al., 2020; Wang et al., 2021). These findings highlight the limitations of climatological
484 baselines and emphasize the necessity of incorporating storm-responsive physical processes.

485 The utilization of the Kullback–Leibler Divergence (KLD) presents notable diagnostic
486 benefits by detecting structural errors such as variance inflation, skewness shifts, and tail
487 distortions, which are not identified by MAE and RMSE (Al-Jarrah et al., 2023; Cover and
488 Thomas, 2005; Freedman and Diaconis, 1981; Kullback and Leibler, 1951; Manimegalai and
489 Bhagyaveni, 2019; Sason and Verdu, 2016). This highlights the importance of adopting
490 distribution-aware metrics as the standard in storm-time validation, particularly in applications
491 where communication is of critical importance.

492 These findings collectively highlight the imperative for the development of hybrid or
493 data-assimilative models that integrate solar-wind coupling, $E \times B$ drifts (Fejer, 2002), and
494 variations in thermospheric composition. These models must be evaluated using both



495 magnitude-based and distribution-sensitive metrics (Adolfs et al., 2022; Astafyeva et al., 2015;
496 Jiang et al., 2024; Luo et al., 2023; Tang et al., 2020; Tsurutani et al., 2008). Such approaches
497 are vital for advancing our ability to accurately model storm-time ionospheric dynamics and
498 enhance the reliability of GNSS-enabled systems.

499 **5 Conclusion**

500 By integrating a conjugate-point framework with distribution-aware validation (KLD),
501 this study provides a novel benchmark for assessing empirical ionospheric models during
502 geomagnetic storms. We utilized empirical ionospheric models as baselines for quiet-time
503 conditions to quantify deviations during storm periods at magnetic conjugate points by
504 employing the Mean Absolute Error (MAE), Root Mean Square Error (RMSE), and
505 distribution-aware Kullback–Leibler Divergence (KLD). The analysis of the storm from August
506 25 to 27, 2018, indicated that the performance of the models was significantly contingent upon
507 the storm phase; all models exhibited degradation during the main phase, with NeQuick2
508 demonstrating superior performance during the recovery phase. Pronounced hemispheric
509 asymmetries were observed; errors were more substantial, and distributions were more distorted
510 at northern mid-latitudes (Haifa), particularly at the equator (Malindi–Djibouti), reflecting
511 genuine geophysical influences such as illumination, winds, and magnetic geometry.

512 The Kullback-Leibler Divergence (KLD) effectively identified structural discrepancies,
513 including variance, skewness, and tails, which remained undetected when the Mean Absolute
514 Error (MAE) or Root Mean Square Error (RMSE) appeared moderate. For instance, during the
515 initial phase of the storm at mid-latitudes, KLD revealed significant distributional distortions
516 in the models (e.g., $KLD=0.30$ for NeQuick2 in Table 1) that were not apparent from the
517 magnitude errors alone ($MAE=1.7$ TECU), highlighting its critical diagnostic value. This
518 underscores the importance of communication-oriented validations. While NeQuick2's solar-
519 flux parameterization offers certain relative advantages, its topside limitations highlight the fact
520 that no empirical model can fully capture storm-time dynamics.

521 Crucially, the robustness checks conducted through the sensitivity analysis
522 (Supplementary Figures S1–S12, Table S2) substantiate that these conclusions are not mere
523 artifacts of parameter selection but rather reflect authentic structural discrepancies.

524 The findings indicate that empirical models should be considered as reference
525 benchmarks rather than predictors. When combined with structure-aware validation, the
526 conjugate-point framework provides a practical methodology for identifying failure modes,



527 prioritizing physical processes for assimilation, and ultimately enhancing the resilience of
528 GNSS and wireless systems under geomagnetic disturbances. This is particularly relevant
529 across magnetically conjugate links, where asymmetries are the most operationally significant.

530 While the current study focused on storms occurring from August 25 to 27, 2018, future
531 research will expand this conjugate-point KLD framework to a statistical ensemble of storms
532 across different seasons and phases of the solar cycle to ascertain the generalizability of the
533 findings.

534 **Code availability**

535 Not applicable. This study used standard Python libraries (NumPy, Pandas, Matplotlib) for data
536 analysis and visualization. No custom code has been archived for public release.

537 **Data availability**

538 Solar wind data are provided by the NASA OMNIWeb database
539 (<https://omniweb.gsfc.nasa.gov>; NASA, 2025). GNSS-derived Total Electron Content (TEC)
540 data were obtained from the International GNSS Service (IGS, <https://igs.org>) and associated
541 global data centers. Processed datasets used in this study, including storm-event subsets for
542 Ambalavao, Djibouti, Haifa, and Malindi stations, are archived at Zenodo and publicly
543 available (DOI: <https://doi.org/10.5281/zenodo.15845821>; Korlaelci et al., 2025).

544 **Interactive computing environment**

545 Not applicable.

546 **Sample availability**

547 Not applicable.

548 **Video supplement**

549 Not applicable.

550 **Supplement link**

551 The link to the supplement will be included by Copernicus, if applicable.

552 **Author contributions**

553 Serhat Korlaelci designed the study, developed the methodology, performed the analysis, and
554 prepared the manuscript. Ramazan Atici contributed to data curation, visualization, and



555 validation. Selçuk Sagir contributed to investigation and writing – review & editing. All authors
556 discussed the results and approved the final manuscript.

557 **Competing interests**

558 The authors declare that they have no conflict of interest.

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