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Dust activities across East Asia and North America have shown decadal variations, mediating radiation budget, air quality, and human health, especially during their peak months of April and May. Using satellite and ground measurements, as along with simulations from a dust emission model, we demonstrate an increase of 3.6% and 30.1% in April dust emissions across East Asia and North America, respectively, during the past four decades, in contrast to a 30.6% and 13.3% decrease during the last two decades. Meanwhile, both regions show a steady increase in May dust emissions by 28.8% and 20.0%, respectively, since the 1980s. Sensitivity experiments attribute both regions' decadal variations in dust emission primarily to surface wind speed changes; whereas vegetation exerts minimum influence on the regional dust emission variations. Furthermore, these decadal variations in dust initiating wind could largely be attributed to regime shifts in extratropical cyclone (EC), including their duration and intensity. These results highlight the changing frequency and duration of strong winds, especially those associated with EC, in shaping the decadal variations of mid-latitude dust emissions.

East Asia and North America are significant dust source regions in the Northern Hemispheric mid-latitudes. The presence and transportation of dust alter the radiation budget (Adebisi et al., 2023;



32 Chen et al., 2013; Huang et al., 2014) and biogeochemical cycles of the marine and terrestrial  
33 ecosystems (Jickells and Moore, 2015; Kong et al., 2022; Jickells et al., 2005). Apart from the  
34 influence on the natural environment, extreme dust activities also impair atmospheric visibility, air  
35 quality, and human health across downwind regions, including populated areas in China and the  
36 United States (Gui et al., 2022; Hashizume et al., 2020; Yang et al., 2015). These environmental  
37 and societal concerns of dust activity peak in April and May across both regions, when vegetation  
38 and snow provide insufficient protection of the dry soil, accompanied by strong near-surface winds  
39 and frequent extratropical cyclones (Aryal and Evans, 2022; Kim et al., 2017; Kurosaki and  
40 Mikami, 2007; Kim et al., 2021). Changes in these atmospheric and land surface factors also shape  
41 the interannual and decadal variations in springtime dustiness across these two mid-latitude  
42 regions, ultimately affecting the regional human well-being.

43

44 Springtime dustiness across both East Asia and North America exhibits substantial interannual-to-  
45 decadal variations, with seemingly opposing trends across decades and divergent driving  
46 mechanisms reported by various studies (Kim et al., 2021; Kurosaki and Mikami, 2007; Xu et al.,  
47 2006). Based on satellite measurements of vegetation greenness and dust aerosol abundance, Wang  
48 et al. (2021) explained the decreased East Asian dustiness in spring by ecological restoration and  
49 resultant vegetation expansion for the period 2001-2020. In contrast, Wu et al. (2022) performed  
50 simulations with a dust emission model and identified surface wind speed weakening as the  
51 dominant driving mechanism for the decreased springtime dust emission across East Asia since  
52 2000s. Apart from that, Song et al. (2021) regarded increasing vegetation and decreasing surface  
53 wind as main contributors of decreasing East Asian dust optical depth during 2009-2019. While  
54 this identified role of wind speed confirmed Tai et al. (2021)'s findings based on chemical transport  
55 model simulations, the latter study, covering the longer period since the 1980s, showed a different  
56 phase of decadal variations in East Asian dust emissions (Tai et al., 2021). Similar debates  
57 regarding the interannual-to-decadal variations in North American dustiness have also persisted.  
58 For example, statistical analysis with ground-based surface fine dust revealed warming and drying  
59 as the main cause of North American springtime dust emission's increase in the early 21<sup>st</sup> century  
60 (Achakulwisut et al., 2017). While according to ground-based dataset in the long-term, Pu and  
61 Ginoux (2018) attributed increased springtime surface fine dust concentrations to the decreasing  
62 trend in precipitation across Southwestern America for the period 1990-2015. Similar spatio-



63 temporal variation was interpreted by Hand et al. (2017) as a result of wind speed, soil moisture  
64 and land cover changes. In contrast, Pu and Ginoux (2017) studied satellite derived dust optical  
65 depth and established the primary relevance between land cover change and Southwestern North  
66 American springtime dust activities for the period 2004-2015. In the context of global warming,  
67 both regions' dust emissions appeared sensitive to vegetation expansion, according to global  
68 climate models and coupled dynamic vegetation-chemical transport model simulations (Li et al.,  
69 2021; Pu and Ginoux, 2017; Zong et al., 2021).

70

71 Despite the rich body of literature on the interannual to decadal variations in springtime mid-  
72 latitude dustiness, little consensus has emerged among these model- and observation-based studies  
73 regarding the direction of change and underlying mechanism. Observational datasets from ground-  
74 and satellite-based measurements provide estimates for atmospheric dust load in the past two to  
75 four decades but lack dust emission records. Therefore, the observed increase or decrease in  
76 atmospheric dust load could be sourced from both local and remote emissions, especially for the  
77 highly transportable mid-latitude dust (Yu et al., 2019b). On the other hand, dust emissions could  
78 be quantified using models, but the credibility of these models, especially their sensitivity to  
79 different influencing factors, should be validated against observation. Moreover, intense dust storm  
80 events that frequently occur in April-May are often modulated by cyclone activity. For instance,  
81 the extreme dust event over East Asia in March 2021 was attributed to both the increased intensity  
82 and frequency of Mongolian cyclones (Liang et al., 2022; Yin et al., 2022). Between 2001-2022,  
83 Mongolian cyclones were reported to contribute 34% to 47% of the total dust emissions from the  
84 Gobi Desert (Mu and Fiedler, 2025). However, quantitative analysis of how the characteristics of  
85 extratropical cyclones, including their occurrence, intensity and size, affect long-term variations  
86 in springtime dust activity across East Asia and North America remains lacking, which limits our  
87 understanding of the drivers behind dust's interannual to decadal variability.

88

89 This study aims to reconcile the interannual to decadal variations in dust emission across East Asia  
90 and North America and quantify the influence of multiple environmental factors. In this work, we  
91 simulate dust emissions across East Asia and North America in April and May from the late 20<sup>th</sup>  
92 century to the early 21<sup>st</sup> century using an observation-validated dust emission model and  
93 subsequently quantify the contribution of meteorology and land surface factors on dust emission



94 changes, including surface wind speed, soil moisture, snow cover fraction, surface temperature,  
95 leaf area index (LAI) from reanalysis and satellite-based observation datasets. Furthermore, this  
96 study integrates cyclone tracking and identification techniques to quantify the impact of  
97 extratropical cyclones on the interannual to decadal variations of both regions' dust activity over  
98 the past four decades. To clarify the reliability of the off-line dust emission model based on Ginoux  
99 et al. (2012), we compare the simulated changes in dust emission with ground-based and satellite-  
100 based dust measurements, such as the global dust Integrated Surface Database (duISD) during  
101 1980-2019, the Interagency Monitoring of Protected Visual Environments (IMPROVE) network  
102 during 1988-2021 and the Moderate-resolution Imaging Spectroradiometer (MODIS) during  
103 2000-2021.

104

## 105 **2. Data and Method**

### 106 **2.1. Ground-based dust measurements**

107 The observed extinction coefficient contributed by dust aerosol ( $\beta$ ,  $km^{-1}$ ) across East Asia ( $35$   
108  $^{\circ}N$ - $50^{\circ}N$ ,  $90^{\circ}E$ - $120^{\circ}E$ ) is provided by global dust Integrated Surface Database (duISD) covering  
109 the period 1980-2019 (Xi, 2021). This dataset compiles about 30,000 stations globally, as collected  
110 by the National Oceanic and Atmospheric Administration (NOAA), and derives dust extinction  
111 coefficient from visibility observations as follows:

$$112 \quad \beta = \frac{3.9}{V} \times f, \quad (1)$$

113 where  $\beta$  is a measure of the extinction coefficient caused by dust particles,  $V$  is the harmonic mean  
114 visibility,  $f$  is the dust frequency (%). 100 and 65 stations provide long-term observations across  
115 East Asia in April and May, respectively, with over two years' valid records during both the late  
116 20<sup>th</sup> century (1980-1999) and the early 21<sup>st</sup> century (2000-2019).

117

118 Parallel to the duISD dataset, the Interagency Monitoring of Protected Visual Environments  
119 (IMPROVE) network has monitored the surface fine dust concentrations ( $\mu g\ m^{-3}$ ) across North  
120 America ( $30^{\circ}N$ - $50^{\circ}N$ ,  $103^{\circ}W$ - $118^{\circ}W$ ) since 1988. The IMPROVE network was originally  
121 designated to support the United States Environmental Protection Agency's Regional Haze Rule  
122 (Hand et al., 2019), and has subsequently been applied to air quality studies, including those on  
123 fine dust variations near the surface (Kim et al., 2021; Pu et al., 2022; Tong et al., 2017). The  
124 IMPROVE dataset provides individual species' contributions to  $PM_{2.5}$  mass and total aerosol



125 extinction twice a week during 1988-2000 and every third day after 2000 in the United States (Pu  
126 et al., 2022). In this work we analyze the surface fine dust concentrations ( $\mu\text{g m}^{-3}$ ) from 1988 to  
127 2021 in April and May.

128

## 129 **2.2. Satellite-based dust measurements**

130 The Moderate-resolution Imaging Spectroradiometer (MODIS) with collection 6.1, level 1  
131 provides the daily dust optical depth (DOD) during 2000-2021 at a spatial resolution of  
132  $0.1^\circ \times 0.1^\circ$ . MODIS DOD is calculated from aerosol optical depth (AOD) and the Ångström  
133 exponent ( $\alpha$ ) as follows:

$$134 \quad \text{DOD} = \text{AOD} \times (0.98 - 0.5089\alpha + 0.051\alpha^2), \quad (2)$$

135 The MODIS instrument is carried by both the Terra (equatorial overpassing at 10:30 a.m.) and  
136 Aqua (equatorial overpassing at 1:30 a.m.) satellites. DOD from MODIS is broadly used in the  
137 study of dust emission and atmospheric loading (Ginoux et al., 2012; Wu et al., 2022; Yu et al.,  
138 2019a; Yu and Ginoux, 2021, 2022).

139

## 140 **2.3. Satellite-based vegetation measurements**

141 The long-term global leaf area index (LAI) is provided by Global Inventory Modeling and  
142 Mapping Studies LAI product (GIMMS LAI4g) (Cao et al., 2023), with a half-month temporal  
143 resolution and a spatial resolution of  $1/12^\circ$  for the period 1982-2020. In this study, we expand the  
144 time range to 1980-2021 by replacing LAI in 1980-1981 with that in 1982 and LAI in 2021 with  
145 that in 2020. The GIMMS LAI4g product used the PKU GIMMS normalized difference vegetation  
146 index product (PKU GIMMS NDVI) and high-quality global Landsat LAI samples to remove the  
147 effects of satellite orbital drift and sensor degradation of Advanced Very High Resolution  
148 Radiometer (AVHRR). In the LAI model simulation, the MODIS Land Cover Type Product  
149 (MCD12Q1, version 6.1) provides the vegetation type reference.

150

## 151 **2.4. Reanalysis data**

152 To investigate the change in dust emissions and the contribution of several environmental variables  
153 in April and May, we analyze the 6-hourly snow cover fraction (%), top layer soil moisture ( $\text{m}^3 \text{m}^{-3}$ )  
154  $^3$ ), land surface temperature (K) and hourly 10-m wind speed ( $\text{m s}^{-1}$ ) from the European Centre for  
155 Medium-Range Weather Forecasts Reanalysis v5-Land (ERA5-LAND, referred to ERA5 hereafter)



156 during 1980-2021. The ERA5-LAND dataset is an enhanced global dataset produced by the  
157 European Centre for Medium-Range Weather Forecasts (ECMWF), with a native resolution of 9  
158 km (Hersbach et al., 2020). The 10-m wind speed from ERA5 can capture the characteristics of  
159 wind to explore wind events both in the hourly and daily scales, compared with station observed  
160 wind speed from Hadley Centre's Integrated Surface Database (HadISD) (Molina et al., 2021).

161

## 162 **2.5. Extratropical cyclone detection and tracking**

163 To analyze regime shifts of extratropical cyclones (ECs) and their contribution to near-surface  
164 strong winds ( $> 6 \text{ m s}^{-1}$ ), we employed the Cyclone TRACKing framework (CyTRACK), an open-  
165 source Python toolbox for cyclone detection and tracking in reanalysis datasets (Pérez-Alarcón et  
166 al., 2024). CyTRACK identifies cyclone centers from mean sea level pressure (MSLP) fields at  
167 each time step and applies threshold-based filtering to track each cyclone. Previous evaluations  
168 have demonstrated that CyTRACK reliably reproduces interannual and seasonal variability, life-  
169 cycle characteristics, and spatial distributions of cyclone tracks when compared with ERA5-based  
170 best-track archives and other cyclone-track datasets.

171

172 In this work, we use the 6-hourly 10-m wind speed ( $\text{m s}^{-1}$ ) and MSLP data in April and May from  
173 ERA5 to identify and track EC during 1980-2021, with a horizontal resolution of  $0.25^\circ$ . Cyclone  
174 centers are mainly identified based on two criteria: (1) surface relative vorticity exceeding  $10^{-5} \text{ s}^{-1}$ ,  
175 a threshold widely applied in EC detection studies (Chen and Di Luca, 2025; Chen et al., 2022;  
176 Priestley et al., 2020), and (2) the central MSLP anomaly being at least 1 hPa lower than the  
177 surrounding grid points (Eichler and Higgins, 2006). Only cyclones with a lifetime longer than 24  
178 hours are retained.

179

180 To quantify the contribution of ECs to strong winds and dust emissions across East Asia and North  
181 America, we determine whether hourly strong-wind events or dust emissions occurred within the  
182 spatial extent of each identified cyclone, defined by the cyclone size radius. The spatial extent of  
183 each cyclone is determined following Schenkel et al. (2017) as the radial distance from the cyclone  
184 center at which the azimuthal-mean 10-m wind speed equals a critical wind speed threshold.  
185 Following previous studies (Pérez-Alarcón et al., 2021; Pérez-Alarcón et al., 2024), we test several  
186 thresholds (2, 4, 6, 8, 10, and  $12 \text{ m s}^{-1}$ ) and adopt  $6 \text{ m s}^{-1}$ , which both aligns with our definition of



187 strong winds and provides the most consistent results. All points within this radius are considered  
188 to be influenced by the cyclone.

189

## 190 **2.6. Offline dust emission model**

191 The quantification of historical, springtime dust emission change across East Asia and North  
192 America is achieved by an offline dust emission model, based on Ginoux et al. (2001) and Ginoux  
193 et al. (2012). Dust emission flux  $F_p$  is calculated as follows:

$$194 \quad F_p = CSu_{max}^2(u_{max} - u_t), \quad (3)$$

195 In April and even occasionally in May, mid-to-high latitude dust source regions in East Asia and  
196 North America are partly covered by snow or frozen, which has a nonnegligible influence on dust  
197 emission (Balkanski et al., 2004; Yin et al., 2022). In this work, we take snow cover fraction, surface  
198 temperature and vegetation cover into consideration, and define the simulated dust emission flux  
199  $F_{p-cover}$  with  $0.1^\circ \times 0.1^\circ$  spatial resolution as follows:

$$200 \quad F_{p-cover} = F_p \times (1 - f_{snow}) \times \exp(-1 \times LAI) \times I, \quad (4)$$

201 where  $f_{snow}$  is daily snow cover fraction,  $LAI$  is daily vegetation cover and  $I$  is the indicator  
202 function of surface temperature (surface temperature  $> 0^\circ\text{C}$ ,  $I = 1$ ; surface temperature  $< 0^\circ\text{C}$ ,  
203  $I = 0$ ).  $C = 1.9 \mu\text{g s}^2 \text{m}^{-5}$  is a dimensional factor,  $S$  is the fraction of dust source (Ginoux et al.,  
204 2010; Ginoux et al., 2012), approximated by the frequency of DOD  $> 0.2$  for the period during  
205 2000-2021 from MODIS in April and May, separately.

206

207  $u_{max}$  is daily maximum surface wind speed in the original model and is substituted with hourly  
208 10-m wind speed in the current study.  $u_t$  is the threshold wind velocity which is calculated as  
209 follows:

$$210 \quad u_t = A \times u_{ref} \sqrt{\frac{\rho_p - \rho_a}{\rho_a} g \Phi_p (1.2 + 0.2 \log_{10} w)} \quad (w < 0.5), \quad (5)$$

211 where  $A = 6$  is a dimensionless parameter,  $u_{ref}$  is a reference threshold wind speed from Pu et al.  
212 (2020).  $\rho_a$  and  $\rho_p$  are the air and particle density,  $g$  is the gravitational acceleration,  $\Phi_p$  is the  
213 particle diameter in five bins:  $0.1-1 \mu\text{m}$ ,  $1-1.8 \mu\text{m}$ ,  $1.8-3.0 \mu\text{m}$ ,  $3.0-6.0 \mu\text{m}$  and  $6.0-20.0 \mu\text{m}$ ,  
214 according to (Kok et al., 2017),  $w$  is the top-layer soil moisture ( $\text{m}^3 \text{m}^{-3}$ ).

215

## 216 **2.7. Sensitivity experiments**



To quantify the contribution of multiple environmental factors on East Asian and North American springtime dust emission changes, we conduct a set of sensitivity experiments that simulate controlled dust emissions. The controlled dust emissions are obtained by individually replacing the concurrent snow cover, temperature, soil moisture, hourly surface wind speed, and LAI during the controlled period with that during the baseline periods. These sensitivity experiments are conducted during the long (1980-2021) and short (2000-2021) terms as follows: in the long-term range, the time subsection is from the late 20<sup>th</sup> century (1980-2000, baseline period) to early 21<sup>st</sup> century (2001-2021, controlled period); in the short-term range, the time subsection is from 2000-2010 (baseline period) to 2011-2021 (controlled period) (Fig. 3). In this study, we use the percentage of dust emission changes between the controlled and baseline simulations to quantify the contribution of each meteorological and biological factor to the decadal changes in dust emission.

229

### 3. Results

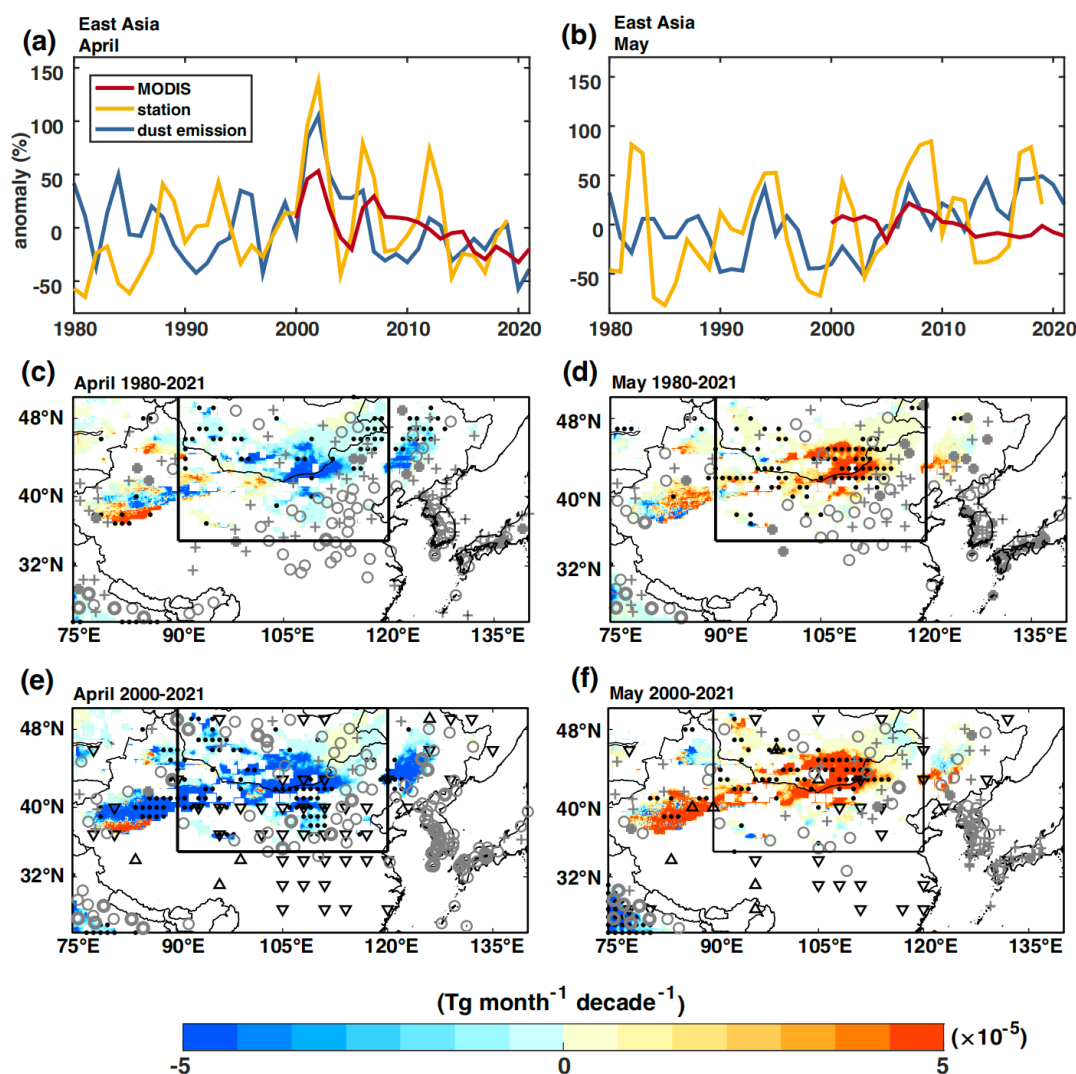
#### 3.1. Decadal variations in East Asian and North American dust emissions

Regional dust emission across East Asia mainly occurs in South Mongolia and North China, and has decreased insignificantly by  $0.252 \text{ Tg month}^{-1} \text{ decade}^{-1}$  in April during the period 1980-2021 (Fig. 1a, c), as also captured by the reduction in dust extinction across the downwind region during 1980-2019 (Fig. 1c). The East Asian dust emission reduction in April is more pronounced after the onset of the 21<sup>st</sup> century, significantly ( $p$ -values  $< 0.05$ , based on the Mann-Kendall trend test) by  $1.82 \text{ Tg month}^{-1} \text{ decade}^{-1}$  or 30.6% per two decades from 2000 to 2021 (Fig. 1a, e). Consistent with this simulated decrease in April dust emission, MODIS DOD exhibits a widespread, significantly ( $p$ -values  $< 0.05$ ) negative trend in April of the early 21<sup>st</sup> century (Fig. 1a, e). Meanwhile, regional dust emission in North America shows a reversed multidecadal trend. North American dust emission has increased significantly ( $p$ -values  $< 0.05$ ) by  $0.261 \text{ Tg month}^{-1} \text{ decade}^{-1}$  or 30.1% per four decades in April during the period 1980-2021, confirmed by the surface fine dust concentrations during 1988-2021, especially in the lower latitudes (Fig. 2a, c). However, this increase is followed by a decrease in the regional total dust emissions by  $-0.388 \text{ Tg month}^{-1} \text{ decade}^{-1}$  or -13.3% per two decades in April for the period 2000-2021, consistently reflected by the satellite-based dust measurement (Fig. 2a, e).





247



248

249 **Figure 1.** Changes of East Asian dustiness in April and May during 1980-2021 and 2000-2021.

250 Time series of DOD anomaly (red lines) from MODIS during 2000-2021, ground-observed

251 extinction coefficient contributed by dust aerosol anomaly (yellow lines) from duISD during 1980-

252 2019 and dust emissions anomaly (blue lines) from off-line simulation model during 1980-2021

253 across East Asia in (a) April and (b) May. Color shading represents the trend of simulated dust

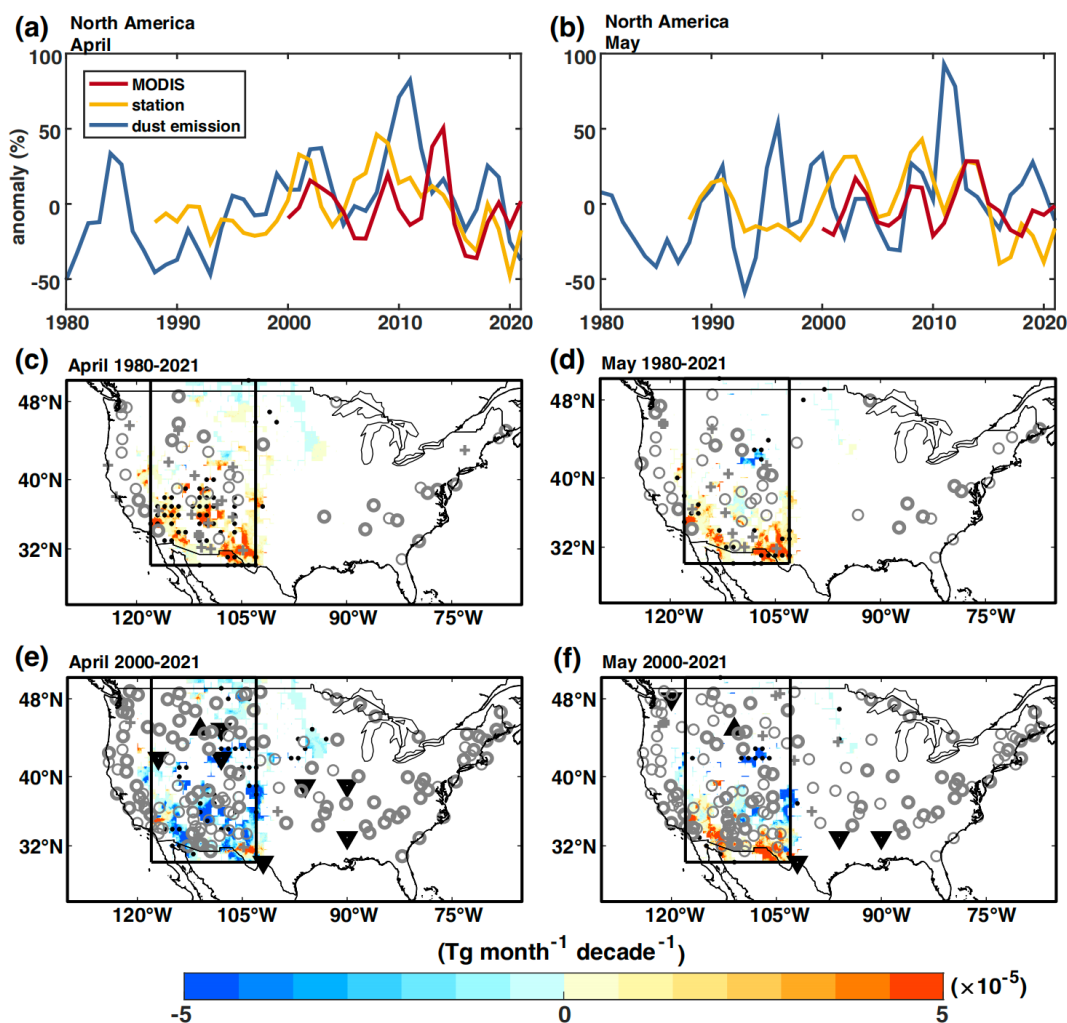
254 emissions (Tg month<sup>-1</sup> decade<sup>-1</sup>) in (c, e) April and (d, f) May for the period (c, d) 1980-2021 and

255 (e, f) 2000-2021. Stippled areas exhibit statistically significant dust emission trends (*p*-values <



0.1, based on the Mann-Kendall trend test). Grey pluses and hollow circles indicate positive and negative, respectively, trends in columnar dust amount during (c, d) 1980-2019 and (e, f) 2000-2019 from duISD. Bold pluses and hollow circles indicate significant ( $p$ -values  $< 0.1$ , based on the Student's  $t$ -test) trends in the corresponding dustiness observation. In (e, f), bold black upward- and downward-pointing triangles indicate significant ( $p$ -values  $< 0.1$ , based on the Mann-Kendall trend test) positive and negative trends in DOD from MODIS in 2000-2021. Boxes denote studied dust source regions across East Asia.

In contrast, both East Asian and North American dust emissions show a steady increase in May during the past four decades. The regional total dust emission is estimated to have increased significantly ( $p$ -values  $< 0.05$ ) by 0.321 and 1.05 Tg month<sup>-1</sup> decade<sup>-1</sup> or 28.8% and 36.7% across East Asia in May for the period 1980-2021 and 2000-2021 (Fig. 1b, d, f). The duISD and MODIS datasets indicates a significant ( $p$ -values  $< 0.1$ ) dust increment across South Mongolia and North China in May for the period 1980-2019 and 2000-2021, respectively (Fig. 1b, d). In North America, the regional total dust emission has increased by 0.226 and 0.171 Tg month<sup>-1</sup> decade<sup>-1</sup> or 20.0% and 15.5% in May during 1980-2021 and 2000-2021 (Fig. 2b, d, f), respectively, consistent with the ground- and satellite-based dustiness observations. Additionally, North American dust emission in May shows a latitudinal variation, with a decrease in dust emissions at higher latitudes and an increase at lower latitudes, particularly in the early 21<sup>st</sup> century (2000-2021) (Fig. 2f).



276

277 **Figure 2.** Changes of North American dustiness in April and May during 1980-2021 and 2000-  
 278 2021. Time series of DOD anomaly (red lines) from MODIS during 2000-2021, ground-observed  
 279 extinction coefficient contributed by dust aerosol anomaly (yellow lines) from IMPROVE during  
 280 1988-2021 and dust emissions anomaly (blue lines) from off-line simulation model during 1980-  
 281 2021 across North America in (a) April and (b) May. Color shading represents the trend of  
 282 simulated dust emissions ( $\text{Tg month}^{-1} \text{ decade}^{-1}$ ) in (c, e) April and (d, f) May for the period (c, d)  
 283 1980-2021 and (e, f) 2000-2021. Stippled areas exhibit statistically significant dust emission trends  
 284 ( $p$ -values  $< 0.1$ , based on the Mann-Kendall trend test). Grey pluses and hollow circles indicate  
 285 positive and negative, respectively, trends in columnar dust amount during (c, d) 1988-2021 and



(e, f) 2000-2021 from IMPROVE. Bold pluses and hollow circles indicate significant ( $p$ -values  $< 0.1$ , based on the Student's  $t$ -test) trends in the corresponding dustiness observation. In (e, f), bold black upward- and downward-pointing triangles indicate significant ( $p$ -values  $< 0.1$ , based on the Mann-Kendall trend test) positive and negative trends in DOD from MODIS in 2000-2021. Boxes denote studied dust source regions across North America.

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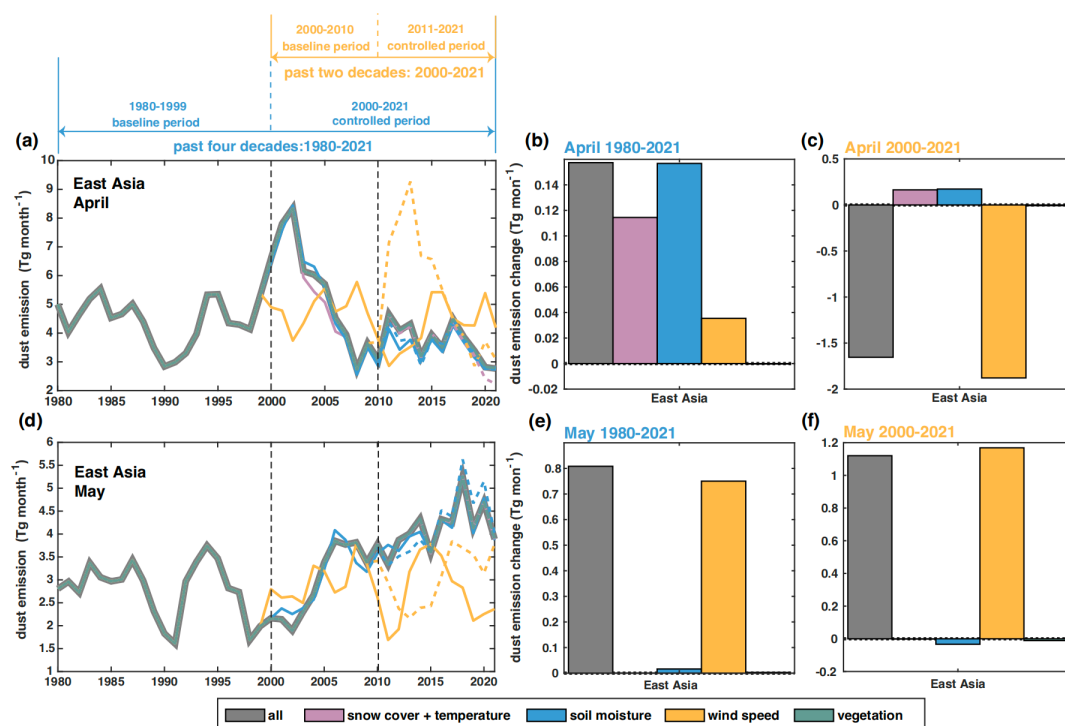
### 292 **3.2. Influencing factors of dust emission changes since the 1980s**

293 The contributions of several environmental variables to the decadal variations in regional dust  
294 emissions are disentangled by the sensitivity experiments. The multidecadal change in East Asian  
295 dust emission in April during 1980-2021 has been mainly driven by soil moisture decrease and  
296 surface warming with reduced snow cover, which individually causes a 3.55% and 2.59% increase  
297 in dust emission (Fig. 3b). Except for this leading role of soil moisture and temperature in dust  
298 emission's multidecadal variation in April across East Asia, surface wind speed emerges as the  
299 main driver of dust emission variations in April and May across both East Asia and North America  
300 (Fig. 3c, e, f, Fig. 4). For example, during 1980-2021, the changes in surface wind speed have  
301 made a positive contribution to dust emission increase by 0.802% and 32.3% across East Asia and  
302 North America, respectively, in April (Fig. 3b, 4b), and a corresponding regional contribution of  
303 26.8% and 13.3% in May (Fig. 3e, 4e). During 2000 to 2021, the surface wind speed has caused a  
304 reduction in dust emissions by -34.8% and -11.3% across East Asia and North America in April  
305 (Fig. 3c, 4c) and an increment by 38.3% and 16.5% in May (Fig. 3f, 4f). On the contrary, changes  
306 in vegetation contribute minimally to dust emission changes over East Asia and North America  
307 during the same periods (Fig. 3, 4), likely due to lack of significant, positive trends in LAI across  
308 both regions in both April and May, especially in the longer term (Fig. 5).

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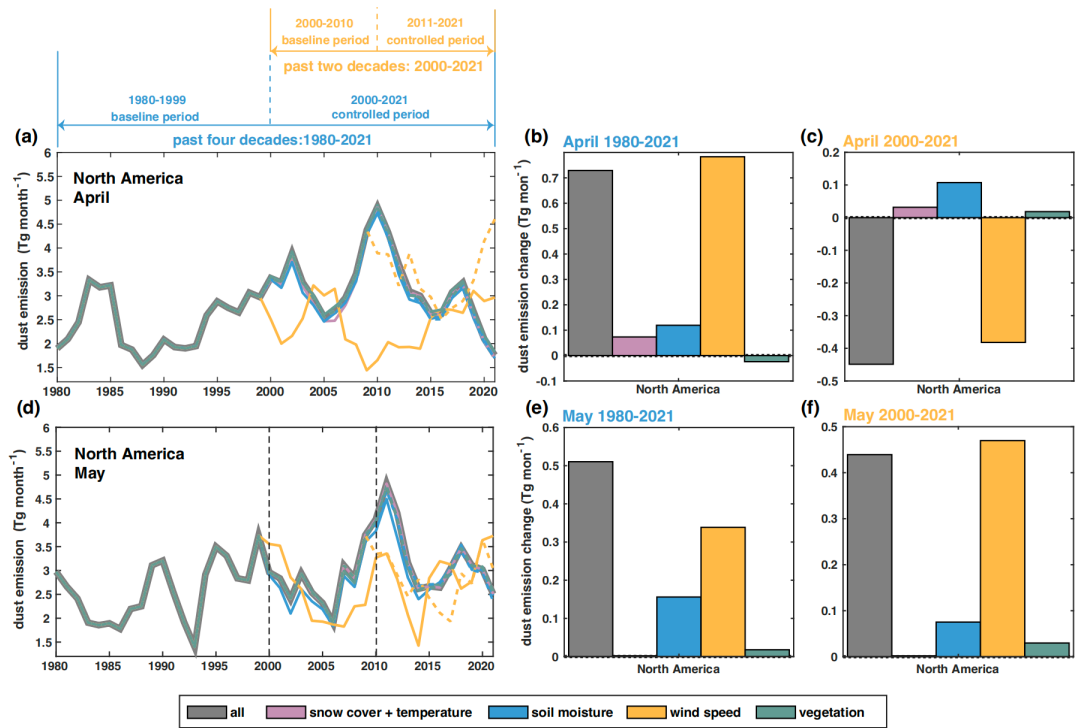
311

312 **Figure 3.** Changes in East Asian dust emission ( $\text{Tg month}^{-1}$ ) and the contribution of several  
313 environmental variables during 1980-2021. Time series of total dust emission (grey) and controlled  
314 dust emission in the sensitivity experiments (colored) across East Asia in (a-c) April and (d-f) May.  
315 The solid, colored lines during 2001-2021 reflect the dust emissions if each environmental variable  
316 repeats their values during 1980-2000; the dashed, colored lines during 2011-2021 reflect the dust  
317 emissions if each environmental variable repeats their values during 2000-2010. (b, c, e, f) Dust  
318 emission change ( $\text{Tg month}^{-1}$ ) and the contribution of each factor (color-coded the same way as in  
319 a, d) across East Asia in (b, c) April and (e, f) May in past four decades (blue font and lines in the  
320 upper of subplot a) and past two decades (yellow font and lines in the upper of subplot a).

321



322



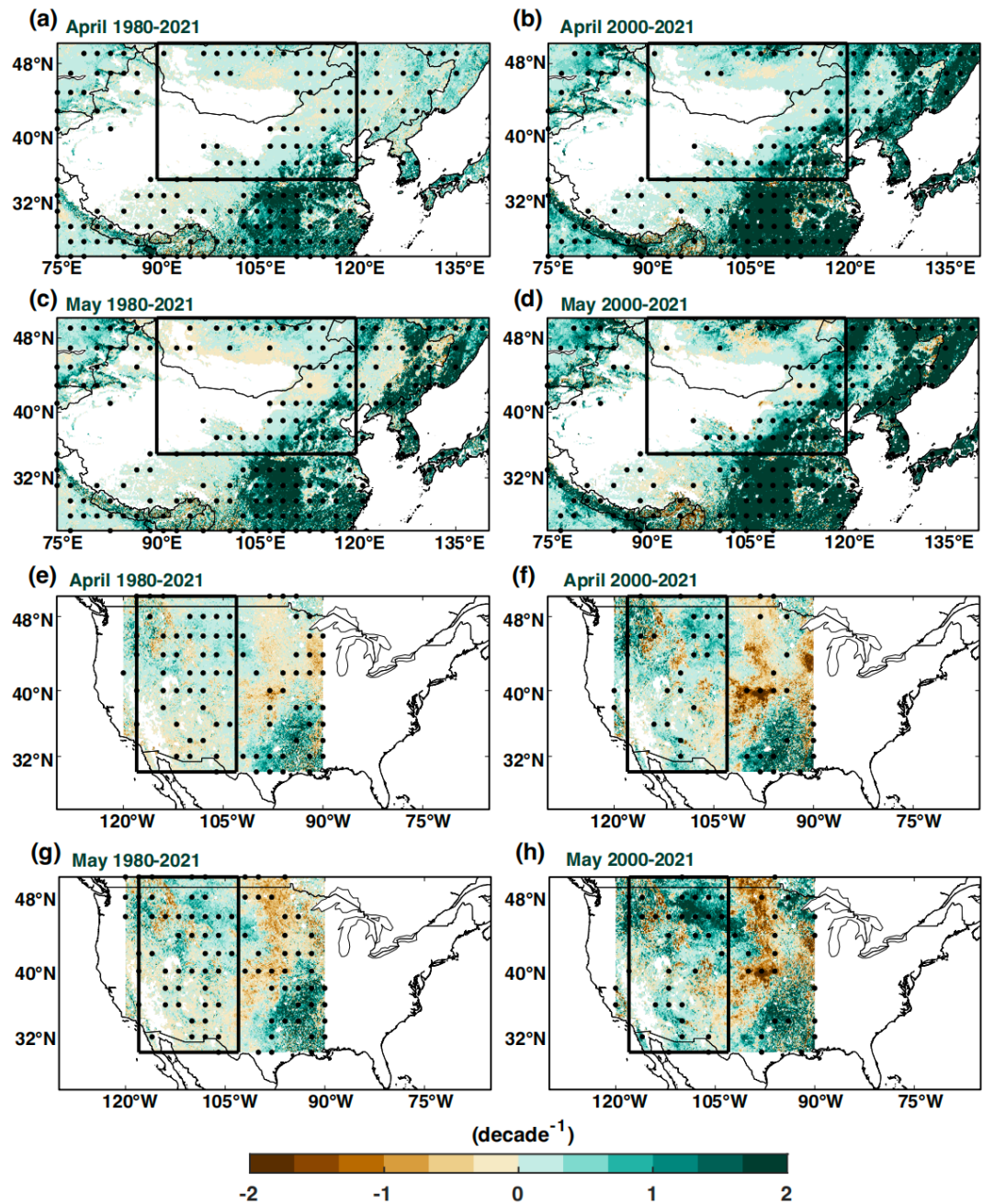
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324 **Figure 4.** Changes in North American dust emission (Tg month<sup>-1</sup>) and the contribution of several  
325 environmental variables during 1980-2021. Figure elements are identical to those in Fig. 3.





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**Figure 5.** Changes of LAI in April and May for the period 1980-2021 and 2000-2021. Trends in (a, b) April and (c, d) May LAI (decade<sup>-1</sup>) across East Asia for the period (a, c) 1980-2021 and



(b, d) 2000-2021. Trends in (e, f) April and (g, h) May daily LAI ( $\text{decade}^{-1}$ ) across North America for the period (e, g) 1980-2021 and (f, h) 2000-2021, respectively. Black dots indicate significant ( $p$ -values  $< 0.1$ , based on the Mann-Kendall trend test) trend in LAI from GIMMS.

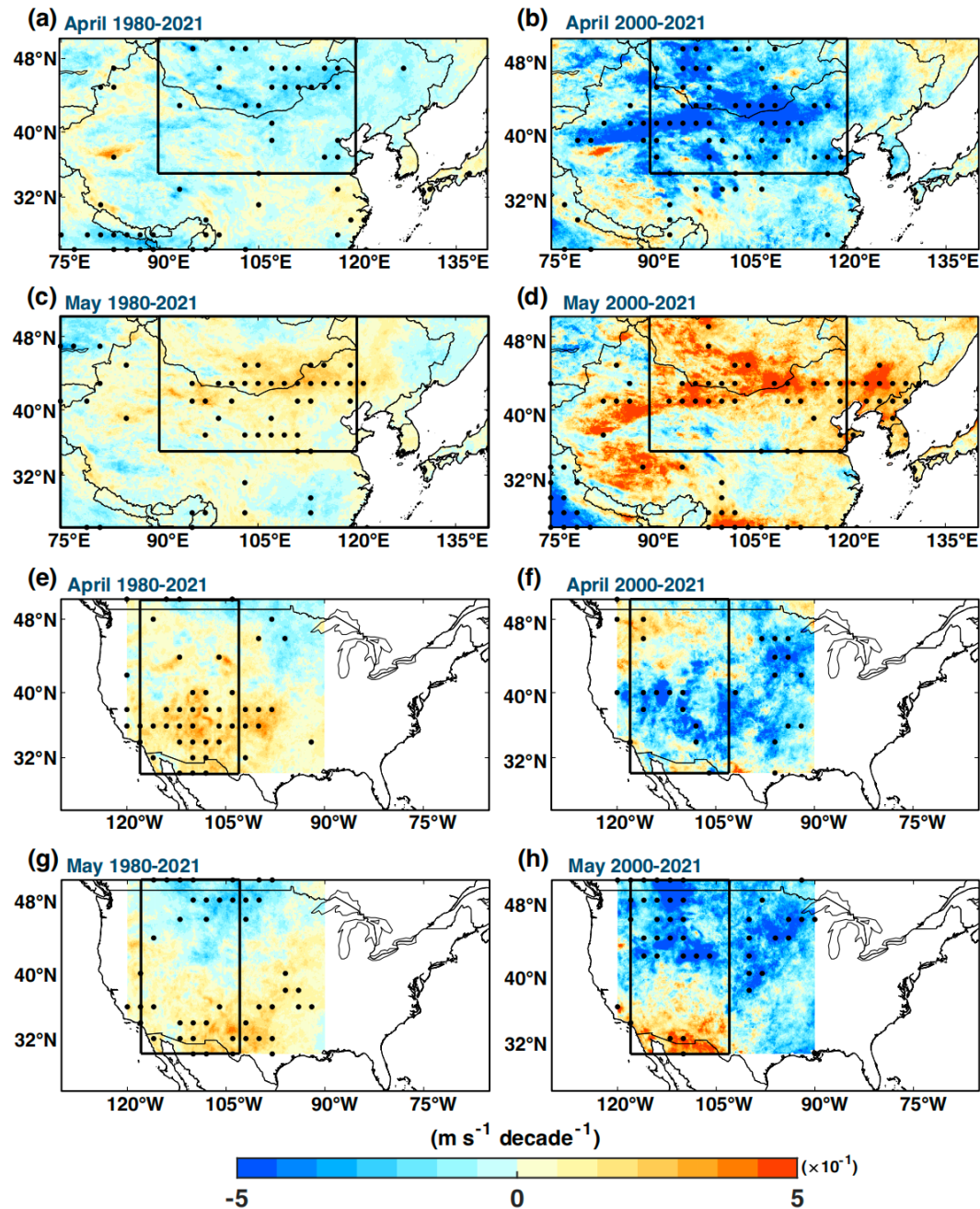
### 3.3 Changing frequency and duration of strong winds

As the dominant influencing factor of East Asian and North American dust emission in mid-to-late spring, near-surface wind speed shows spatio-temporally in-phase variations with dust emission. Spatially, daily maximum wind speed (Fig. 6) exhibits similar patterns of change with those in dust emissions across both regions during both the shorter and longer periods (Fig. 1, 2). Temporally, the positive contribution of wind speed to dust emission across North America in April during 1980-2021 (Fig. 4b) and across both regions in May during the entire period 1980-2021 (Fig. 3e, f, 4e, f) is a consequence of the higher occurrence of longer-lasting strong winds (Fig. 7). For example, the occurrence of strong winds with duration between 150 and 260 hours and frequency between 18 and 34 events per month has increased significantly across North America in April (Fig. 7h), leading to the positive trend of daily maximum wind speed across North America in April from the late 20<sup>th</sup> century (1980-2000) to the early 21<sup>st</sup> century (2001-2021) (Fig. 4). Similarly, strong winds that last for over 150 hours, typically associated with synoptic weather systems, occur more frequently in May across both regions during the recent decades (Fig. 7), resulting in the multidecadal, steady increase in dust emission from both regions in May (Fig. 1, 2, 4).





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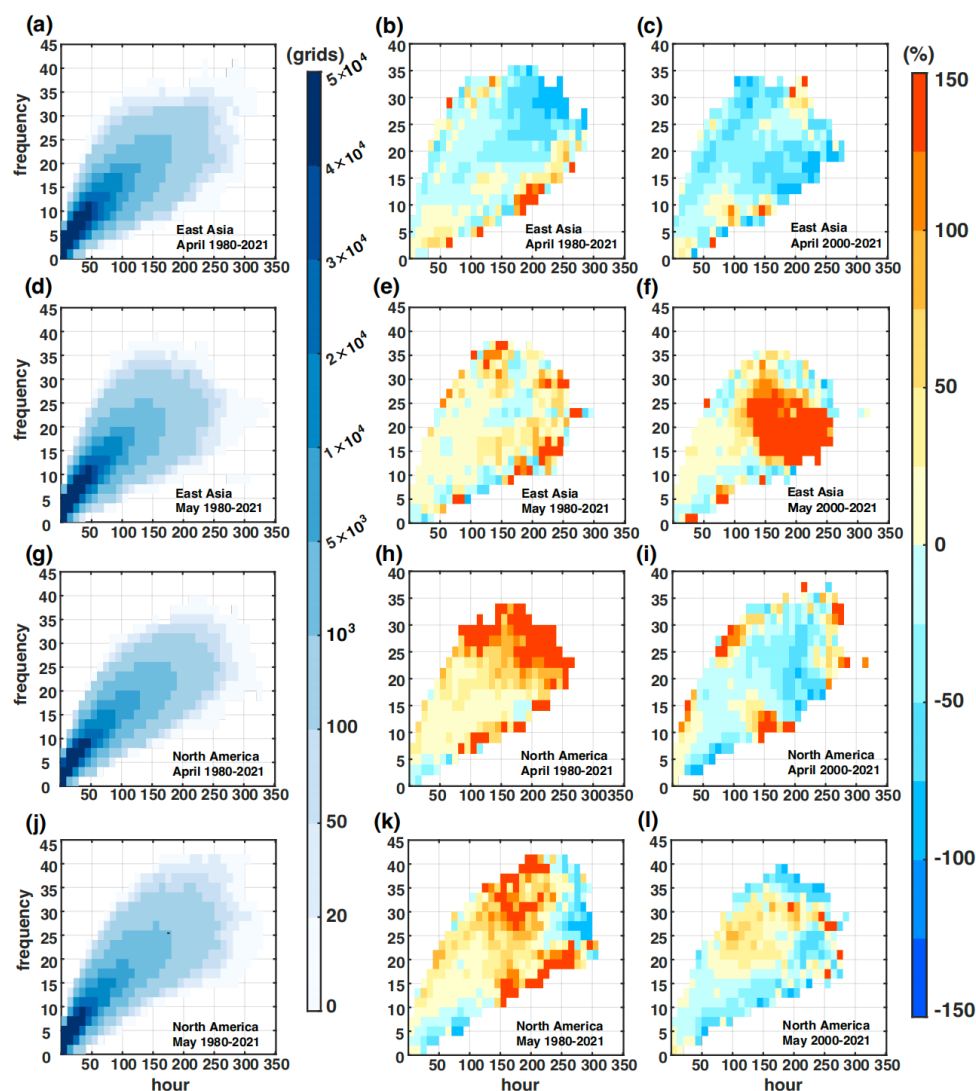
355

**Figure 6. Changes of daily maximum wind speed in April and May for the period 1980-2021 and 2000-2021. Trends in (a, b) April and (c, d) May daily maximum wind speed (m s<sup>-1</sup> decade<sup>-1</sup>)**



356 across East Asia for the period (a, c) 1980-2021 and (b, d) 2000-2021. Trends in (e, f) April and (g, h) May daily maximum wind speed ( $\text{m s}^{-1} \text{ decade}^{-1}$ ) across North America for the period (e, g) 1980-2021 and (f, h) 2000-2021, respectively. Black dots indicate significant ( $p$ -values  $< 0.1$ , based on the Mann-Kendall trend test) trend in daily maximum wind speed.

360



361

362 **Figure 7. Changes of strong wind across East Asia and North America in April and May.**

363 Joint probability distribution (joint PDF) of the monthly total frequency and duration of strong

364 wind ( $> 6 \text{ m s}^{-1}$ ) from ERA5 hourly 10-m wind speed data across the dust-emitting pixels in (a, d)



365 East Asia and (g, j) in North America in (a, g) April and (d, j) May during 1980-2021. Change rate  
366 (%) of the joint PDF of the frequency (events per month) and duration (hours per month) of strong  
367 winds ( $> 6 \text{ m s}^{-1}$ ) to baseline periods from ERA5 hourly 10-m wind speed data across the dust-  
368 emitting pixels in (b, c, e, f) East Asia in (b, c) April and (e, f) May for the period (b, e) 1980-2021,  
369 (c, f) 2000-2021 and in (h, i, k, l) North America in (h, i) April and (k, l) May for the period (h, k)  
370 1980-2021, (i, l) 2000-2021.

371

### 372 **3.4 Regime shifts of extratropical cyclones**

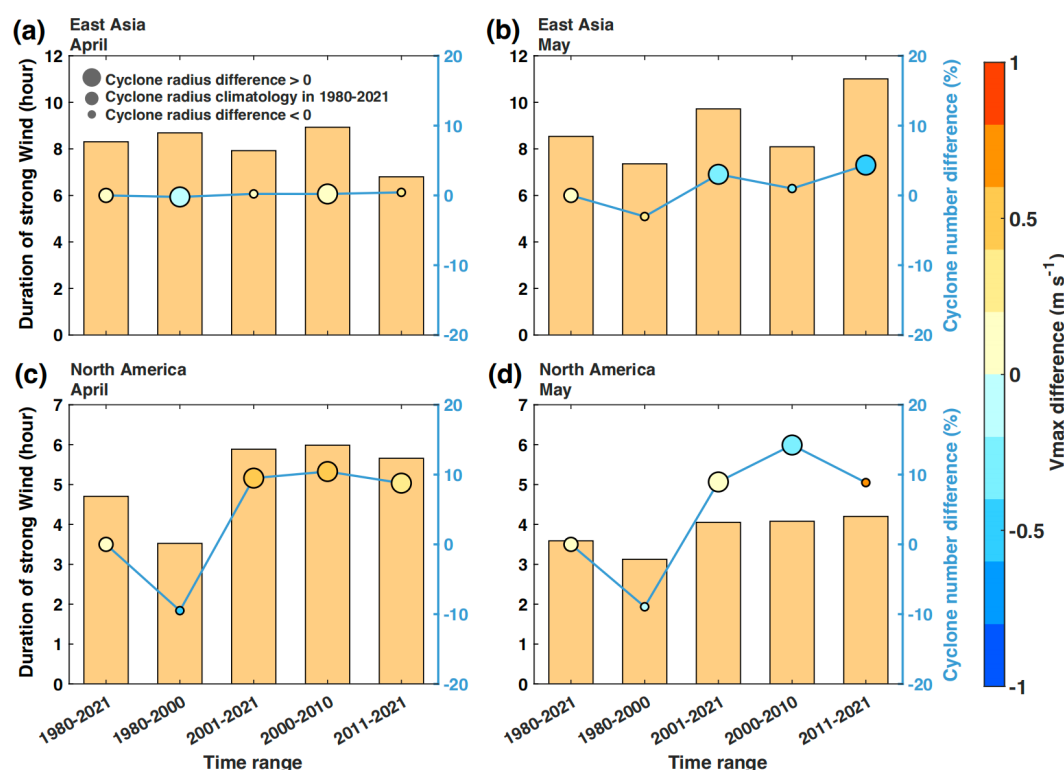
373 Extratropical cyclones exert a significant influence on near-surface strong winds during April and  
374 May, which in turn affect dust emission processes. In East Asia, the annual average occurrence and  
375 maximum wind speed ( $V_{\text{max}}$ ) of extratropical cyclone, duration of cyclone-driven strong winds,  
376 and cyclone-induced dust emissions showed no substantial changes in April from 1980 to 2021  
377 (Fig. 8a, 9a). However, all four factors exhibited distinct inter-decadal variability during 2000-  
378 2021 (Fig. 8a, 9a). This pattern aligns with the currently identified, minimal overall contribution  
379 of wind to dust emission changes in East Asia in April from 1980 to 2021 (Fig. 3a, b, 7b), and  
380 explains the decline in dust emissions associated with reduced strong winds from 2000 to 2021  
381 (Fig. 3a, c, 7c). In May, the increase in both the duration of strong winds and cyclone frequency  
382 shaped the frequency and duration of strong winds' change pattern in East Asia between 1980-  
383 2021 and 2000-2021 (Fig. 7e, f, 8b), thereby driving the inter-decadal variability in dust emissions  
384 (Fig. 3d, e, f, 9b).

385

386 The increased North American dust emissions in April from 1980 to 2021, driven by longer-lasting  
387 and more frequent strong winds (Fig. 4b, 7h), also can be attributed to changes in extratropical  
388 cyclones. During this period, both the number of extratropical cyclones and their maximum wind  
389 speed increased, resulting in longer durations of strong winds (Fig. 8c). In contrast, from 2000 to  
390 2021, the decrease in the number of cyclones and  $V_{\text{max}}$ , along with shorter periods of cyclone-  
391 caused strong winds, leads to the decline in wind-driven dust emissions in April, which dominate  
392 the inter-decadal variability of dust emissions (Fig. 4c, 7i, 8c, 9c). In May, the changes in strong  
393 winds during 1980-2021 can also be explained by variations in extratropical cyclones and their  
394 associated strong wind durations. More frequent and stronger cyclones cause extended periods of  
395 strong winds, leading to higher North American dust emissions (Fig. 4e, 7k, 8d, 9d). However, for



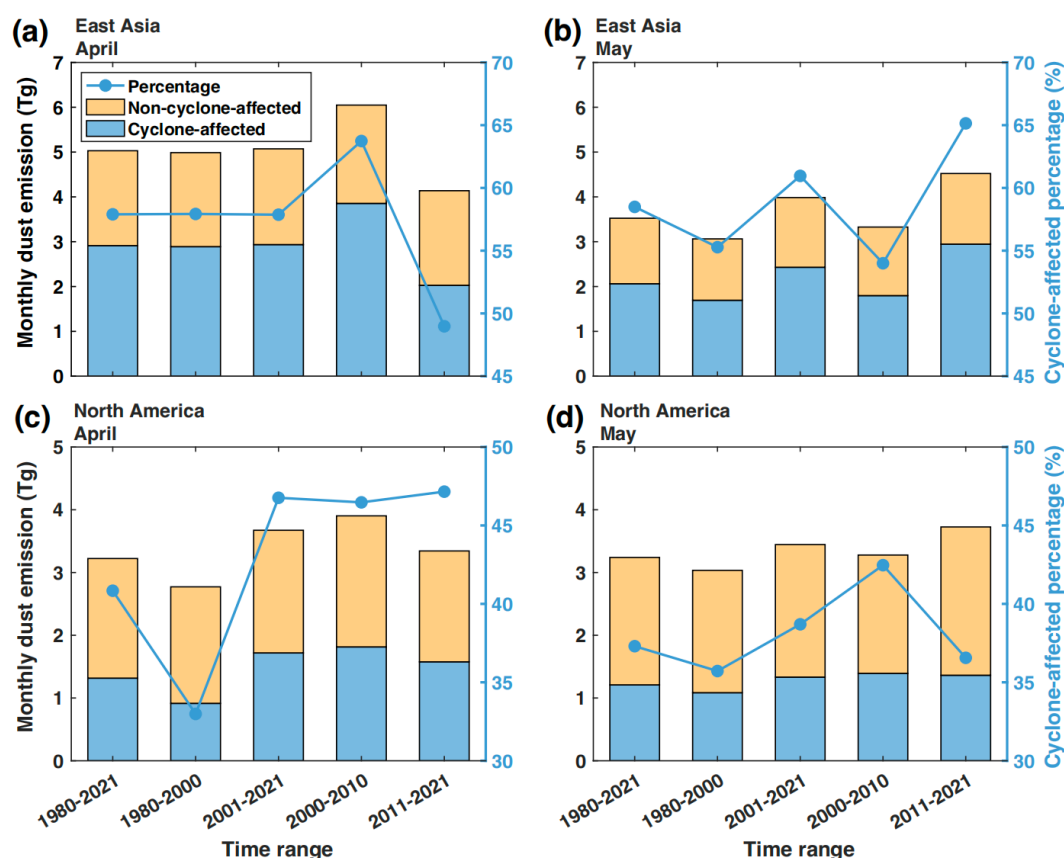
the period in 2000-2021, although the duration of cyclone-caused strong winds increases slightly (Fig. 8d), the percentage of dust emissions driven by extratropical cyclones have weakened (Fig. 9d). This may be explained by the reduced number and radius of cyclones (Fig. 8d), which lead to a relative decrease in the area affected by cyclone-induced dust activity; meanwhile, within the cyclone-affected regions, stronger winds enhanced dust emission intensity. During the period in 2000-2021, cyclone-induced dust emissions remain nearly constant, whereas dust emissions from non-cyclone strong wind events increased (Fig. 9d), likely due to enhanced smaller-scale or weaker-magnitude disturbances.



**Figure 8.** Regime shifts in extratropical cyclones across East Asia and North America during April and May across different time periods from 1980 to 2021. The monthly average duration of strong winds (hours; yellow bars) caused by extratropical cyclones in (a, b) East Asia and (c, d) North America in (a, c) April and (b, d) May, for the time periods: 1980-2021, 1980-2000, 2001-2021, 2000-2010, and 2011-2021, with reference to the left y-axis. The blue solid line with markers



411 represents the difference in the monthly average number of cyclones during these periods  
412 compared to the monthly average cyclone count during the whole period 1980-2021,  
413 corresponding to the right blue y-axis. Marker color shows the deviation of monthly mean Vmax  
414 ( $\text{m s}^{-1}$ ) from the 1980-2021 climatology, and marker size reflects the cyclone radius difference  
415 relative to the 1980-2021 mean.  
416



417  
418 **Figure 9.** Regime shifts in extratropical cyclone-caused dust emissions (Tg) across East Asia and  
419 North America in April and May across different time periods from 1980 to 2021. The annual  
420 average dust emissions (Tg) along the passage of extratropical cyclones (blue bars) and those  
421 unaffected by them (yellow bars) are shown for (a, b) East Asia and (c, d) North America in (a, c)  
422 April and (b, d) May, for the time periods: 1980-2021, 1980-2000, 2001-2021, 2000-2010, and  
423 2011-2021, with reference to the left y-axis. The blue solid line represents the percentage of dust



emissions affected by extratropical cyclones (%) over these periods, corresponding to the right y-axis.

#### 4. Discussion and conclusion

Based on a suite of multi-source observational datasets and a dust emission model, we characterize the decadal variability of mid-to-late springtime dust emissions across East Asia and North America, which are primarily regulated by changes in surface wind speed and extratropical cyclone activity during the recent decades. During the past four decades, the East Asian and North American drylands exhibit a 4% and 30% increase in April dust emissions and a 29% and 20% increase in May. During the past two decades, these two regions show a 31% and 13% decrease in April dust emissions and a 37% and 16% increase in May. Our results highlight the dominant role of surface wind speed in shaping decadal variations of dust emissions, while the frequency and intensity of extratropical cyclones exert substantial influence on wind speed variability. Collectively, these two factors constitute the primary drivers of regional total dust emission changes across East Asia and North America in the late 20<sup>th</sup> century and early 21<sup>st</sup> century. Overall, our study provides a clearer understanding of the decadal-scale variability of mid-to-late springtime dust emissions across East Asia and North America, and underscores the primary roles of both surface wind speed and extratropical cyclones in modulating dust emission changes.

In this study, we demonstrate the leading influence of surface wind speed on decadal changes of dust emission. Wind speed regimes, especially the frequency, duration, and intensity of strong wind events, reflect the signal of both climate variability and climate change. Extratropical cyclones exert a dominant influence on near-surface strong winds, which in turn drive dust emissions. The identified decadal changes in near-surface wind speed, along with the changing duration of strong wind events, can be largely attributed to regime shifts in extratropical cyclones (Fig. 7, 8, 9). In addition, these changes may also be interpreted as part of (1) the response in mid-latitude storm track processes to global warming (Shaw et al., 2016), (2) regional climate oscillations associated with large scale modes of climate oscillation, such as El Niño-Southern Oscillation (ENSO), North Atlantic Oscillation (NAO), Pacific Decadal Oscillation (PDO), Arctic Oscillation (AO) etc. (Yin et al., 2022), and (3) global surface wind stilling up to 2010 and subsequent recovery attributed to internal climate variability (Zeng et al., 2019; Wohland et al., 2021).





455

456 Compared with wind speed, land surface changes seem secondary in shaping the decadal variations  
457 in dust emission, except the slight dust emission increase across East Asia in April since the 1980s  
458 in response to surface warming, reduced snow cover, and soil moisture. Beyond reflecting the  
459 mixed signals from climate variability, land surface factors directly respond to climate change. For  
460 example, studies on vegetation phenology have reported an earlier greening trend across Northern  
461 Hemispheric mid-latitudes in response to early-spring warming and CO<sub>2</sub> fertilization (Fan et al.,  
462 2014; Piao et al., 2019). But vegetation's increasingly stronger inhibition on dust emission appears  
463 unable to offset the influence of surface wind speed for the period 1980-2021 (Fig. 5). Furthermore,  
464 future changes in vegetation cover highly depend on the competing trajectories of surface  
465 temperature and soil moisture, and their role in dust emission remains uncertain (Ding et al., 2020).

466

467 The uncertainty of our study mainly comes from the limitation of observational datasets and dust  
468 emission model. First, station observations generally lack temporal continuity and spatial coverage  
469 especially in dust source regions; therefore, the long-term trend in dustiness from stations likely  
470 miss the complete spatio-temporal pattern. Second, higher albedo of cloud and land surface, in the  
471 presence of thick clouds and snow, respectively, brings challenge to satellite aerosol retrieval  
472 algorithms, preventing a more accurate quantification of dust concentration or emission solely  
473 based on satellite data. Third, although the simulation from off-line dust emission model generally  
474 matches observed spatio-temporal variations, this parameterization inevitably underrepresents  
475 actual physical processes, similar to all dust emission models currently being used, especially the  
476 interaction between environmental variables. For example, we estimate the area of unvegetated,  
477 wind-erosive regions within each grid by  $\exp(-1 \times LAI)$  (Pu and Ginoux, 2017). This  
478 parameterization, however, omits the influence of vegetation height and canopy structure on near-  
479 surface wind profile and eventually the frictional wind speed that is directly responsible for dust  
480 emission. This uncertainty in dust emission modeling will be quantified and reduced upon an  
481 expanded collection of observable data, e.g. meter-resolution vegetation structure, spatio-  
482 temporally resolved near-surface wind speed profiles, in conjunction with satellite measurement  
483 of dust aerosol abundance with finer spatio-temporal resolutions.

484

485 **Data Availability Statement**



486 Data used in this study are all publicly available, including: the MODIS Deep Blue aerosol  
487 products acquired from the Level-1 and Atmosphere Archive and Distribution System (LAADS)  
488 Distributed Active Archive Center (DAAC) (<https://doi.org/10.5194/amt-6-949-2013>); the ERA-5  
489 hourly climate data provided by European Centre for Medium-Range Weather Forecasts (ECMWF)  
490 (<https://doi.org/10.5194/essd-13-4349-2021>); the GIMMS leaf area index at a half-month temporal  
491 resolution acquired from Cao et al. (2023); the global dust Integrated Surface Database (duISD)  
492 acquired from Xi (2021); the Interagency Monitoring of Protected Visual Environments  
493 (IMPROVE) network is available for download (<http://vista.cira.colostate.edu/improve>).

494

#### 495 **Code Availability**

496 The code to carry out the current analyses is available from the corresponding authors upon request.

497

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502

#### 503 **Author contributions**

504 YY conceived the study. YW and YY performed the analysis and wrote the initial draft. All authors  
505 contributed to the data analysis and manuscript editing.

506

#### 507 **Competing interests**

508 The authors declare that they have no conflict of interest.

509

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