



# Machine learning-driven characterization and prescription of aerosol optical properties for atmospheric models

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## Abstract

Accurate modeling of aerosol optical properties is critical to simulate aerosol radiative effects. However, uncertainties regarding the simulation aerosol intensive optical properties are still significant. Therefore, the use of observations to constrain aerosol optical properties in models has been indicated as an option. Also, explicit computations of optical properties are still too costly for operational models, which make observational-based prescriptions a convenient solution. We developed a observational-based prescription of aerosol optical properties driven by machine-learning techniques that can be applied in models. The Iberian Peninsula (IP) was taken as the reference domain, and the aerosol products from the AERONET sites across the IP as the main dataset. First, clustering was applied to define the typical aerosol optical regimes affecting the IP atmosphere. Five typical regimes were identified. Two of them were dominated by coarse mode, which were associated with Saharan dust. One was found to be close to pure dust, while the other indicated a mixed scenario of dust and pollution. Two of the non-dust regimes, strongly and moderately absorbing, were found to be associated with smoke. The remaining non-dust regime, with not a clear association, occurs mostly in the eastern portion of the IP. Afterward, using aerosol-type columnar mass density from MERRA-2, a model was trained as predictor of the optical regimes using the Random Forest method. The model was tested under distinct aerosol scenarios. Predictions' accuracy ranged from 60 to 75%, depending on the regime, while presenting an average accuracy of 70%.

**Keywords:** Aerosol Optical Properties, AERONET, MERRA-2, Machine-Learning, Random Forest



## 36 1. Introduction

37 Aerosol particles' importance in the Earth's climate system is undisputed. Via the scattering  
38 and absorption of terrestrial and solar radiation, aerosol particles are direct players in the  
39 planetary energy budgets (Kim and Ramanathan 2008, IPCC, 2021, Li et al., 2022). However,  
40 this role is permeated by high complexity and significant uncertainty (Spencer et al. 2019,  
41 IPCC, 2021, Li et al., 2022). The uncertainties and challenges in accurately representing  
42 aerosol particles' processes in climate, weather, and environmental models arise from  
43 various limitations. For instance, focusing on aspects related to the direct interaction with  
44 radiation, limitations in the current global observational system to address information such  
45 as spectral complex refractive index and size distribution, two critical microphysical  
46 variables to the characterization of particle absorption and scattering (Samset et al. 2018, Li  
47 et al., 2022), are still a relevant source of uncertainty.

48 The lack of geographical representativity of the traditional libraries of aerosol optical and  
49 microphysical properties (Shettle and Fenn, 1979, Koepke et al., 1997, Hess et al., 1998) has  
50 been central in the aerosol optical properties uncertainty debate. Another critical aspect is  
51 the characterization of the state of the mixture of the aerosol particles in the model's aerosol  
52 modules (Samset et al. 2018, Sand et al., 2021). Given the complex dynamic of aerosol particle  
53 emission, transport, and removal in the atmosphere, numerical modelling of the state of the  
54 mixture and the resultant complex refractive index and size distribution is widely recognized  
55 as one of the most important sources of uncertainty in addressing aerosol particles' radiative  
56 forcing (Sand et al., 2021). According to Sand et al. (2021) aerosol absorption is poorly  
57 constrained, and the current climate models present a large range in the quantification of the  
58 main absorbing aerosol species (black carbon (BC), organic aerosols (OA), and mineral dust).  
59 Brown et al. (2021) findings indicate that biomass-burning aerosols in most climate models  
60 are too absorbing mainly due to treatments of aerosol mixing state. Saharan dust, a critical  
61 component of the global aerosol system, has been found to absorb less solar radiation than  
62 models estimate (Adebisi et al., 2023), and the primary cause pointed out is the models  
63 overestimate of the dust imaginary refractive index. Absorption is not the only issue facing  
64 aerosol particle representation in climate models, the relative contribution of fine and coarse  
65 mode particles is also a challenge. For instance, Adebisi et al. (2023) also found models  
66 underestimating large dust particles when representing North African dust plumes.

67 Observation-constrained models have been recommended to mitigate models' current  
68 difficulty in fully simulating aerosol properties and processes accurately (Samset et al. 2018,  
69 Proske et al., 2024). In addition to the uncertainty aspects, explicit simulation of aerosol  
70 compositions and microphysical, followed by explicit computations of intensive optical  
71 properties, is still too expensive computationally for operational models, which also makes  
72 observational-based prescriptions a convenient solution. Zhong et al. (2022) used  
73 relationships from an ensemble of aerosol models and satellite observations to identify the  
74 primary source of uncertainty in aerosol modelling results. Their study pointed out the  
75 incorrect lifetimes and the underestimation of mass extinction coefficients as the most  
76 critical drivers of bias in aerosol simulations. As the largest, time and device consistent  
77 observational network, capable of constraining multiple aerosol intensive microphysical and  
78 optical properties, the AErosol RObotic NETwork (AERONET) has been used worldwide to



79 constrain models and satellite algorithms (Omar et al., 2005, Li et al., 2010, Levy et al., 2010,  
80 Rosario et al., 2013, Russel et al., 2014, Chen et al., 2023). Chen et al. (2023) developed an  
81 aerosol optical module with observation-constrained Black Carbon properties to improve  
82 aerosol absorption simulation. Their sensitivity simulations show a reduction of 18%–69%  
83 in the biases of aerosol single-scattering co-albedo when compared with global observations  
84 from AERONET. Li et al. (2010) used AERONET retrievals to evaluate and improve the  
85 performance of a GCM aerosol optical module. They found their GCM to simulate flatter  
86 Aerosol Optical Depth (AOD) spectral dependence, indicating an Angstrom Exponent (AE)  
87 biased to low values, which suggests that the aerosol sizes simulated were too large. After  
88 adjusting the aerosol's size based on AERONET retrievals the agreement between simulated  
89 and observed AOD improved for all aerosol regimes, but especially for smoke and dust  
90 scenarios. Rosario et al. (2013) used a set of spectral optical models developed from  
91 AERONET sky retrievals over distinct biomes combined with the concept of anisotropic areas  
92 of influence of the AERONET sites (Hoelzemann et al., 2009) to constrain smoke aerosol  
93 radiative effect modelling during South America biomass burning. By doing so, they were  
94 able to capture the effect of the regional variability of smoke optical properties (absorption  
95 and size related) on the surface solar irradiance related to the biomes' distinct nature of  
96 smoke.

97 Global and regional cluster analysis of AERONET long-term retrievals of aerosol properties  
98 has proved valuable to classify observations in terms of aerosol optical regimes, providing  
99 means to qualitative constraints on aerosol properties (Omar et al., 2005, Levy et al., 2007,  
100 Russell et al., 2014, Li et al., 2019, Fan et al., 2020, Zhou et al., 2023). In these studies, the  
101 number of the identified typical aerosol optical regimes varied from 4 to 10, numbers that  
102 were expected to likely represent either global or regional major aerosol scenarios  
103 variability, according to each study focus. In their study, Zhou et al. (2023) found regional  
104 aerosol regime classifications to perform better than global classifications when applied to  
105 simulate AOD during pollution episodes and in different seasons in Beijing, China. They found  
106 a large difference between the strongly and moderately absorbing aerosol regimes in the  
107 global and regional clustering results. Two major sources of differences between their global  
108 and regional clustering for China were aerosol optical regimes dominated by dust and smoke  
109 particles. Compared to China, Zhou et al. (2023) pointed out that smoke and dust-dominated  
110 optical regimes are more common globally. Their result suggests that regional classification  
111 better captures typical aerosol optical regimes influencing a specific domain and, therefore,  
112 with potential to improve observation-constrained simulations of aerosol radiative forcing.

113 Focusing on the Iberian Peninsula (IP), this study sought to characterize the typical aerosol  
114 optical regimes driving the variability of aerosol-intensive properties over the peninsula,  
115 aiming to constrain aerosol optical properties prescription in atmospheric models using a  
116 novel approach based on machine-learning approach. IP is a region affected by a highly  
117 dynamic and complex set of aerosol mixing, including natural and anthropogenic particles  
118 (Cachorro et al., 2016, Gomez-Amo et al., 2017). Natural sources include marine aerosols  
119 from the Atlantic Ocean and Mediterranean Sea, mineral dust from North Africa, and  
120 eventually, wildfire emissions. Major anthropogenic sources are urban-industrial,  
121 particularly in more densely populated regions, and biomass burning driven by human  
122 activities, especially in the north and central Portugal and eastern and north of Spain.



123 Regional column-integrated optical properties are highly sensitive to the mixing of this  
124 diversity of aerosol-types, in particular to dust and smoke mixing (Gomez-Amo et al., 2017).

125 The manuscript is organized as follows: Section 2 includes a brief overview of the Iberian  
126 Peninsula, focusing on the main atmospheric circulation features and major aerosol particle  
127 sources affecting the region, followed by the description of the dataset and methods adopted  
128 to identify, characterize and prescribe the identified aerosol typical regimes. Results and  
129 discussions are presented in Section 3. First, the identified aerosol optical regimes and their  
130 major features are described and contextualized. Subsequently, the results of the novel  
131 machine-learning approach to prescribing the optical regimes are discussed and evaluated.  
132 Finally, the main findings of our study are highlighted in the conclusion section.

133

## 134 **2. Study Region, Data and Methods**

135

### 136 **2.1 Study region**

137 The Iberian Peninsula (**Figure 1**), comprising Spain and Portugal, exhibits diverse climate  
138 conditions due to its complex topography and proximity to the Atlantic Ocean, the  
139 Mediterranean Sea and North Africa. The wind circulation over the peninsula is shaped by its  
140 location between the Atlantic Ocean and the Mediterranean Sea, diverse topography, and  
141 interactions between regional and global atmospheric patterns, leading to complex wind  
142 circulations that significantly influence the region's climate. This results in distinct climate  
143 zones, from arid deserts to lush green forests. The Mediterranean climate spans most of  
144 Spain, including the eastern and southern coastal regions and central Portugal, featuring hot  
145 and dry summers, especially inland. Winters are mild, rarely dropping below 10°C in coastal  
146 areas. Most precipitation, often rain, occurs in autumn and winter, leading to dry summers  
147 that increase wildfire risks. Wildfires regularly occur in the IP region fueled by extreme  
148 weather conditions, abnormal high temperature records combined with strong, dry winds  
149 (Asfaw et al., 2022, Ermitão et al., 2023). Under these scenarios the entire region can be  
150 affected by smoke plumes that often shape the entire region's optical properties (Elias et al,  
151 2004, Gomez-Amo et al., 2017). But wildfires are more frequent in the north and central  
152 region of Portugal and the north and eastern portion of Spain (Ermitão et al., 2023, Alvares  
153 et al., 2024). Oceanic climate is typical in northern coastal regions of Spain, such as Galicia,  
154 Asturias, and the Basque Country, and parts of northern Portugal. The Atlantic Ocean  
155 influences mild temperatures year-round, with minimal seasonal variation and abundant,  
156 evenly distributed rainfall. Annual precipitation can exceed 1,000 mm, with frequent cloud  
157 cover and high humidity, especially in winter. The Continental climate of the central plateau  
158 (Meseta Central) and the Ebro Valley features extreme temperature variations, with hot  
159 summer, highs often above 35°C, and winter below freezing. The central regions have less  
160 precipitation than the coastal areas, with a semi-arid climate in some parts. Most rainfall  
161 occurs in spring and autumn. Arid and Semi-Arid Climates are found in Southeastern Spain,  
162 especially in Murcia and Almería, and parts of the Ebro Valley. These areas receive very low  
163 rainfall, often less than 300 mm annually, leading to desert-like conditions like those in the  
164 Tabernas Desert. Summers are extremely hot, while winters are mild. Southern Spain,

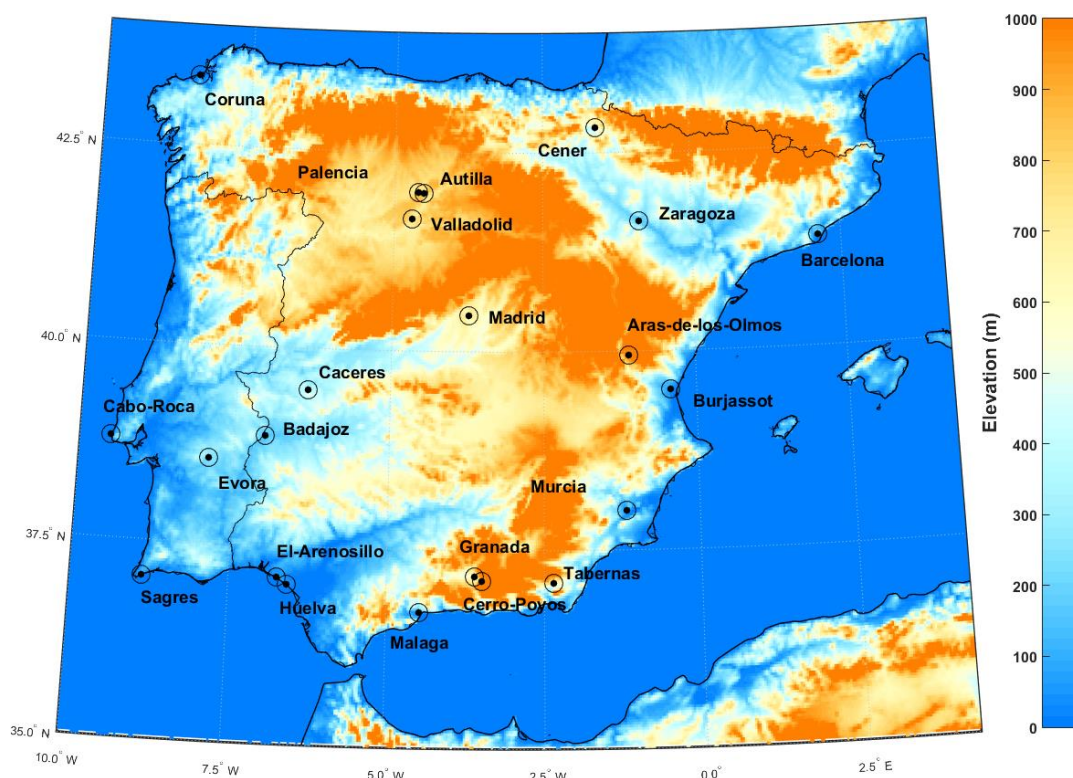


165 especially the Andalusian region, can be affected by hot and dry winds from the Sahara,  
166 causing heat waves and dust storms.

167 The occurrence of Saharan dust events on Iberian Peninsula usually peaks in March and June,  
168 with a marked minimum in April and lowest occurrence in winter according to Cachorro et  
169 al. (2016). Depending on the synoptic conditions and circulation patterns, dust transport can  
170 affect the entire peninsula (Toledano et al., 2007). The prevailing westerlies, blowing from  
171 west to east, are the dominant wind pattern over the Iberian Peninsula. These winds are most  
172 prominent in the mid-latitudes, including the Iberian Peninsula. More pronounced in the  
173 northern region during autumn and winter, these winds bring moist air from the Atlantic,  
174 increasing precipitation in Galicia, the Basque Country, and north Portugal. While they also  
175 affect central and southern areas, their impact is moderated by the peninsula's topography  
176 and other wind systems. The northeast trade winds affect the southern and western coasts  
177 of Portugal and southwestern Spain, creating a mild and dry climate, especially in summer.  
178 In contrast, Mediterranean winds affect the eastern and southeastern coasts. Additionally,  
179 the Iberian Thermal Low, resulting from intense heating of the Iberian interior, creates a low-  
180 pressure area that draws air from the Atlantic and Mediterranean shores, leading to  
181 converging wind patterns. This circulation pattern enhances sea breeze penetration and  
182 moderates coastal temperatures. Southern Spain is influenced by Sahara winds, as said these  
183 dry winds often carry dust, increasing temperature and reducing air quality. Calima is a type  
184 of wind that occurs when Saharan dust reaches the peninsula, especially in summer, causing  
185 hazy skies, a reddish tint, and low visibility. These winds are linked to high-pressure systems  
186 over North Africa and low-pressure systems over the western Mediterranean.

187 The wind circulation over the Iberian Peninsula is a dynamic and complex system shaped by  
188 global atmospheric patterns, regional geography, and local topography. The interaction of  
189 prevailing westerlies, trade winds, Mediterranean breezes, and local wind systems creates a  
190 diverse wind regime that affects the peninsula's climate. Understanding these patterns is  
191 essential for weather prediction, agriculture management, and tackling environmental  
192 challenges. According to Cachorro et al. (2016), these complex and contrasting influences of  
193 air masses from the Atlantic Ocean, Mediterranean Sea, European continent, and North Africa  
194 lead to a large spatio-temporal variability in aerosol properties, types, and mixing processes  
195 over the Iberian Peninsula. This makes the peninsula a challenging region for online  
196 modeling of aerosol microphysical properties and mixing state, therefore an interesting  
197 region to evaluate observation-constrained approaches.





**Fig. 1:** AERONET sites locations displayed on top of the Iberian Peninsula topography.

## 2.2 AERONET aerosol inversion product

AERONET is a global ground-based network of sun photometers mainly aimed at characterizing columnar aerosol particles properties (Holben et al., 1998). From the direct Sun attenuation measurements, AERONET algorithms derive spectral Aerosol Optical Depth ( $AOD_\lambda$ ) at the wavelengths 0.34, 0.38, 0.44, 0.50, 0.67, 0.87, 0.94, and 1.02  $\mu\text{m}$ . From the spectral dependency of AOD at these wavelengths, AERONET provides Angstrom Exponent (AE), a parameter sensitive to the aerosol particle size distribution (Eck et al., 1999). But for the present study, AERONET also provides several other intensive properties that depend not on the amount but on the nature of the aerosol, related to particle size, shape, and composition, from sky radiance measurements at the wavelengths 0.44, 0.67, 0.87, and 1.02  $\mu\text{m}$  (Sinyuk et al., 2020). These intensive properties include microphysical parameters, such as refractive indices ( $n+ik$ ) and volume size distribution, and also optical parameters like Single Scattering Albedo (SSA), asymmetry parameter (ASY), Lidar Ratio (LR), Linear Depolarization Ratio (LDR), Angstrom Exponent among others (Holben et al., 1998, Dubovik et al. 2002). Given the dependency of these intensive properties on the aerosol type and mixture state, it is possible to characterize the aerosol scenarios over a specific AERONET site in terms of their nature and sources (Eck et al., 1999; Dubovik et al., 2002). Therefore,



218 with a well-distributed regional network of AERONET' sun photometers, as that covering  
219 Iberian Peninsula, one can characterize the spatial dynamic of aerosol types and mixture  
220 state influencing the regional aerosol regimes.

221 Three key aspects of aerosol nature have been widely used to link aerosol regimes with  
222 particle emission sources. These aspects are absorption efficiency, size distribution and  
223 shape (Dubovik et al., 2002). For instance, combustion-based sources, including biomass and  
224 fossil fuel burning, produce aerosol dominated by fine mode particles, and absorption ranges  
225 from moderate to strong, depending on biomass burning nature, fossil fuel and ageing  
226 processes. In contrast, natural sources, such as deserts and marine environments, produce  
227 aerosols dominated by coarse-mode particles. Marine aerosol particles are characterized by  
228 very low absorption, while dust aerosol can exhibit high absorption, mainly in the UV and VIS  
229 bands (Smirnov et al., 2002; Dubovik et al. 2002). Furthermore, the irregular shape of dust  
230 particles is a key factor that differentiates them from other aerosol types. This distinctive  
231 feature is captured by AERONET retrievals of the LDR (Shin et al., 2018). Source attribution  
232 provides valuable insights into the typical intensive optical properties affecting the  
233 atmospheric column of a site resulting from complex aerosol state mixtures. This  
234 understanding is crucial as it addresses a major challenge that current aerosol modules in  
235 climate models face. Reproducing climatological aerosol-intensive properties scenarios over  
236 specific regions has been a major goal of atmospheric models. In addition to evaluating  
237 aerosol modules in atmospheric models, AERONET's optical properties typical regimes,  
238 which can be expressed as spectral aerosol optical models (Omar et al., 2005; Levy et al.,  
239 2007; Rosario et al., 2013; Zhou et al, 2023), are valuable for simulating aerosol direct  
240 radiative effects in environmental models (Rosario et al., 2013, Li et al., 2019). This approach  
241 is especially beneficial when/where high computational capacity is unavailable and explicit  
242 aerosol modules are not feasible.

243 Aiming to identify a representative set of typical aerosol regimes that affect the Iberian  
244 Peninsula, we applied cluster analysis methods (described in Sec. 2.4) to the AERONET  
245 dataset, taking advantage of the extensive coverage of AERONET sites across the region.  
246 **Table 1** presents a set of intensive properties provided by AERONET that was used to  
247 identify typical aerosol scenarios in the Iberian Peninsula atmospheric column. The variables  
248 displayed cover all the three previously mentioned aspects, absorption efficiency, size  
249 distribution and shape, which are expected to characterize the distinct nature of aerosol  
250 types and mixture anticipated in the study region. We selected only AERONET sites that  
251 operated for at least two years and that have sky radiance inversion available with the  
252 highest quality level 2.0. Some selected sites are still operational, while others have been  
253 discontinued. **Figure 1** illustrates the geographical distribution of the chosen sites. Our  
254 selection encompasses various landscapes of the Iberian Peninsula, from coastal plains  
255 regions (Coruña, Sagres, Burjassot) to highland plateau in the interior (Madrid, Valladolid,  
256 Aras-de-los-Olmos) and lowland valleys (Zaragoza, Murcia). Regarding external air mass  
257 influence, sites in the southern border of IP are typically the first to experience the transport  
258 of dusty air mass from North Africa, with locations such as El- Arenosillo, Huelva, Malaga,  
259 Sagres affected. The eastern sites (Barcelona, Burjassot, Murcia) are expected to be strongly  
260 influenced by the Mediterranean air masses. Western and northern sites (Cabo da Roca,  
261 Coruna, Sagres) are directly under the influence of air mass from the Atlantic Ocean.



262 Additionally, Portugal countryside (Evora) and Spain eastern sites (Badajoz, Caceres) are  
263 located in regions that very often experience biomass burning during the dry season  
264 (Ermitão et al., 2023, Silva et al., 2023, Hammed e tal., 2024, Alvares et al., 2024).

265

266 **Table 1:** List of AERONET sky inversions intensive properties variables used in clustering  
267 process.

Variables	Abbreviation
Refractive Index - Real Part	RI <sub>Real</sub> (440), RI <sub>Real</sub> (670), RI <sub>Real</sub> (870), RI <sub>Real</sub> (1020)
Refractive Index - Imaginary part	RI <sub>Imag</sub> (440), RI <sub>Imag</sub> (670), RI <sub>Imag</sub> (870), RI <sub>Imag</sub> (1020)
Single Scattering Albedo	SSA(440), SSA(670), SSA(870), SSA(1020)
Asymmetry Parameter	ASY(440), SSA(670), SSA(870), SSA(1020)
Linear Depolarization ratio	LDR(440), LDR(670), LDR(870), LDR(1020)
Lidar Ratio	LR(440), LR(670), LR(870), LR(1020)
Fine and Coarse modes Volume median radius	VMR-F,VMR-C
Standard deviation from volume median radius, for Fine and Coarse modes	STD-F, STD-C
Fine and Coarse modes Effective radius	Reff-F, Reff-C

268

## 269 2.3 Merra-2 Aerosol Diagnostic Product

270 The MERRA-2 (Modern-Era Retrospective Analysis for Research and Applications, Version 2)  
271 Aerosol Diagnostic Product (ADP) is a comprehensive dataset provided by NASA that offers  
272 global information about atmospheric aerosols (Gelaro et al., 2017, Buchart\_Marchand et al.,  
273 2017). MERRA-2 combines observational data with numerical models(reanalysis project) to  
274 create a detailed long-term record of atmospheric dynamics and composition from 1980 to  
275 the present. Among other variables, the MERRA-2 ADP product offers a long-term view of  
276 aerosol mass distribution by types and the related optical properties (Bucharth\_Marchand et  
277 al., 2017). Its extended temporal coverage allows analysis of aerosol trends, such as those  
278 related to changes in atmospheric composition due to human activity and the impact on  
279 climate. Key features of the MERRA-2 ADP include aerosol microphysical and optical  
280 properties such as optical depth, mass concentration, and size distribution. These properties  
281 are crucial for understanding aerosol loading and composition in the atmosphere and their  
282 role in the Earth's radiation budget and climate system. A key aspect of MERRA-2 APD for  
283 this study is that it provides aerosol-type column mass density, our target variable as a  
284 predictor of aerosol optical model regime. The MERRA-2 APD includes diagnostics for the  
285 aerosol types considered in most chemistry transport models: Dust (DT), Black-Carbon (BC),





286 Organic Carbon (OC), Sea-Salt (SS) and Sulfate (SF). The aerosol-type diagnostics variables  
287 cover mass concentration at specific levels and integrated in the entire atmospheric column,  
288 as well as columnar optical properties, such as extinction, scattering and absorption optical  
289 depth. From these extensive aerosol-driven optical properties, it is possible to derive several  
290 MERRA-2 ADP intensive optical properties, such as Single Scattering Albedo (SSA).

291 Given that the aerosol optical properties retrieved from each AERONET site are influenced  
292 by the mixture of different aerosol types present in the local atmospheric column, it is  
293 reasonable to assume that the impact of each aerosol type on the column's intensive optical  
294 properties is primarily determined by its concentration. Based on this premise, we propose  
295 a machine-learning approach that utilizes the aerosol-type column mass density predicted  
296 by chemistry transport models to help us define the spatial distribution of the optical model  
297 developed through cluster analysis of AERONET data. A description of the method presented  
298 in this study, exploring MERRA-2 products, can be found in subsection 2.5.

299

## 300 **2.4 Optical models development: Cluster Analysis**

301 Cluster analysis has been extensively used to develop aerosol optical models based on  
302 AERONET sky inversion products (Omar et al., 2005, Levy et al., 2007, Russel et al., 2014). The  
303 underlying principle is that AERONET instantaneous retrievals can be grouped into a certain  
304 number of clusters, each representing different categories of aerosol regimes. These studies  
305 have explored mainly the K-means clustering method, one of the most popular unsupervised  
306 machine learning algorithms for partitioning a dataset into a pre-defined number of clusters.  
307 However, specifying the number of clusters in advance poses a significant challenge for the  
308 K-means method. Fortunately, there are techniques available that minimize the subjectivity  
309 involved in this pre-definition. In our study, we adopted the Elbow method (Shi et al., 2021),  
310 the most widely used method for determining the optimal number of clusters (k) in a K-  
311 Means clustering algorithm. It examines the relationship between the number of clusters and  
312 the within-cluster sum of squares (WCSS), which measures the variance within each cluster  
313 (**Eq. 1**)

$$314 \quad \text{WCSS} = \text{minimize}(\sum_{k=1}^k W(C_k)) \quad (1)$$

315 where  $C_k$  is the  $k$ th cluster and  $W(C_k)$  is the within-cluster variation. The total within-cluster  
316 sum of squares (WCSS) measures the compactness of the clustering, and one wants it to be  
317 as small as possible. We ran our clustering algorithm with  $k$  varying from 2 to 10 clusters.  
318 For each  $k$ , we calculated the total within-cluster sum of squares (WCSS). The  $k$  results  
319 against WCSS were displayed in a plot and the optimal number of clusters were defined based  
320 on the location ( $k$ ) of the bend (elbow) in the plot.

321

## 322 **2.5 Optical models spatial prescription: Random Forest Technique**

323 We propose a machine-learning approach that utilizes the well-known random forests  
324 supervised algorithm (Breiman, 2001) to spatially represent the aerosol optical models  
325 defined by the cluster analysis for each AERONET site (described in section 2.4). The



326 implemented method was tested using aerosol column mass density data from MERRA-2  
327 (**Table 2**) to establish the spatial distribution of the optical regime defined by the clusters  
328 average. This approach is also suitable for chemistry transport models.

329 MERRA-2 time series of column mass density for each aerosol type (DT, BC, OC, SS, SF) over  
330 each AERONET site were collocated with the network inversion products used to derive the  
331 clusters representing the distinct aerosol regimes over the Iberian Peninsula (described in  
332 section 2.4). Each AERONET instantaneous aerosol microphysical and optical properties  
333 inversion retrieval (Sinyuk et al., 2020) was connected to the corresponding cluster to which  
334 it belonged. Likewise, each instantaneous aerosol microphysical and optical properties  
335 inversion retrieval was also connected to the closest in-time combination of MERRA-2 data  
336 of aerosol-type column mass density (DT, BC, OC, SS, SF). With this, we built a set of data fitted  
337 to a training process, wherein the occurrence of each cluster (optical model) is related to a  
338 particular mixture (combination) of aerosol types from MERRA-2 over each selected  
339 AERONET site. The nature of our problem is classification, with the combination of the  
340 aerosol-type columnar mass density, trying to predict which cluster of intensive optical  
341 properties is more suitable for that particular combination.

342 Therefore, the first step was to split the data into training (70%) and test (30%). The  
343 algorithm uses training data to learn the relationship between the combination of aerosol-  
344 types columnar mass density and the target, which are the developed clusters from  
345 AERONET aerosol-intensive properties. The training was done using the Random Forest  
346 Classification algorithm (RandomForestClassifier) from the Python package Scikit-Learn  
347 (Abraham et al. 2014). The Random Forest classifier's hyperparameters were optimized  
348 using RandomizedSearchCV, a stochastic method of parameter space exploration. The  
349 parameter space included the number of decision trees (n\_estimators: 50–500) and the  
350 maximum depth of trees (max\_depth: 1–20). The process used stratified k-fold cross-  
351 validation to ensure representative sampling across aerosol regime classes. This  
352 optimization method addressed the issues of class imbalance and aerosol regime  
353 classification in atmospheric measurements. The random search methodology was used to  
354 find parameter combinations inside the parameter space without the processing demands of  
355 grid search. Cross-validated performance indicators were used to select the final  
356 configuration in order to reduce overfitting and ensure consistent performance across  
357 aerosol regimes. The confusion Matrix was used to visualize the performance of the models,  
358 and we also calculate the following indicators: Accuracy, Precision and Recall and F1  
359 score. Accuracy represents the number of correctly classified data instances over the total,  
360 it checks the predictions against the actual values in the test set and returns the percentage  
361 of times the model got right.

362 Precision and recall are two critical metrics for evaluating the performance of a classification  
363 model. Precision is the proportion of true positives among all the predicted positive cases  
364 (true and false), meaning it measures the accuracy of positive predictions (**Eq. 2**). Recall is  
365 the proportion of true positives among all actual positive cases (true and false), meaning it  
366 measures the model's ability to identify positive cases (**Eq. 3**). The F1 score, the harmonic  
367 mean of a model's precision and recall, takes both precision and recall into account and  
368 provides a more balanced measure of a model's performance (**Eq. 4**). The F1 score is set to



be a value between 0 and 1, indicating, respectively, poor precision and recall and high precision and recall, which is ideal.

$$\text{Precision} = \text{True positive} / (\text{True positive} + \text{False positive}) - (2)$$

$$\text{Recall} = \text{True positive} / (\text{True positive} + \text{False negative}) - (3)$$

$$F1 = 2 \times (\text{Precision} \times \text{Recall}) / (\text{Recall} + \text{Precision}) - (4)$$

**Table 2:** Predictor variables from Merra-2 (aerosol-type column mass density) used in the machine learning process to prescribe the aerosol optical regime (optical model).

Variables	Abbreviation	Unity	Spatial resolution
Dust column mass density	DUCMASS	kg/m <sup>2</sup>	0.5° × 0.625°
Black carbon column mass density	BCCMASS	kg/m <sup>2</sup>	0.5° × 0.625°
Organic carbon column mass density	OCCMASS	kg/m <sup>2</sup>	0.5° × 0.625°
SO <sub>2</sub> column mass density	SO2CMASS	kg/m <sup>2</sup>	0.5° × 0.625°
SO <sub>4</sub> column mass density	SO4CMASS	kg/m <sup>2</sup>	0.5° × 0.625°
Sea salt column mass density	SSCMASS	kg/m <sup>2</sup>	0.5° × 0.625°

### 3. Results

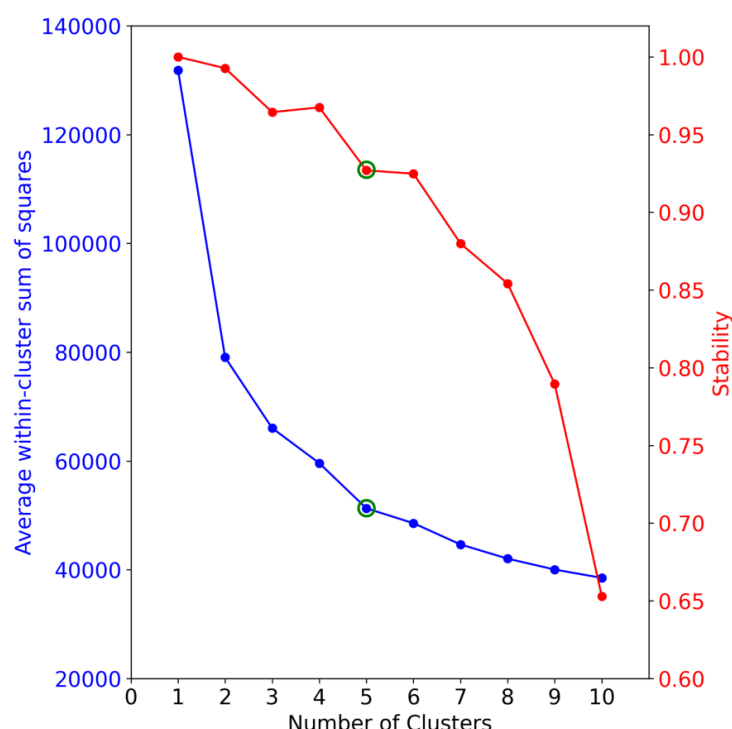
The results section is divided into three subsections. The first one presents the results of identifying the typical aerosol optical regimes affecting the Iberian Peninsula using cluster analysis. The second subsection discusses the results and the performance of spatial prescription of these typical aerosol regimes by applying machine learning (Random Forest) to the columnar density of MERRA-2 aerosol components. Finally, case studies applying the method developed are presented and discussed.

#### 3.1 Cluster Analysis: Optical models development

The number of clusters ( $k$ ) selected to characterize the typical optical aerosol regimes over the Iberian Peninsula was defined based on the Elbow method (**Figure 2**), which indicated five as the optimal clusters number to capture the aerosol regime variability. We also evaluated from the Elbow method that there is a sharp bending at  $k=2$ , which we associated with a clustering separation between aerosol regimes strongly dominated by coarse mode, dust regimes, and regimes dominated by fine mode, non-dust regimes. However, to cover more specific regimes within these two macro-regimes (dust regimes vs non-dust regimes) a higher  $k$  is required, and  $k=5$  reveals to be the second sharpest bending. Cluster stability as



a function of the number of clusters was also evaluated as a complementary analysis. The stability for  $k=5$  is above the 90% threshold, similar to  $k=6$ , a number after which stability sharply decreases. Therefore, combining the Elbow method and stability reinforced  $k=5$  as an optimal cluster number to capture the typical aerosol scenarios over the Iberian Peninsula, reducing the subjectivity usually associated with the K-mean clustering method.



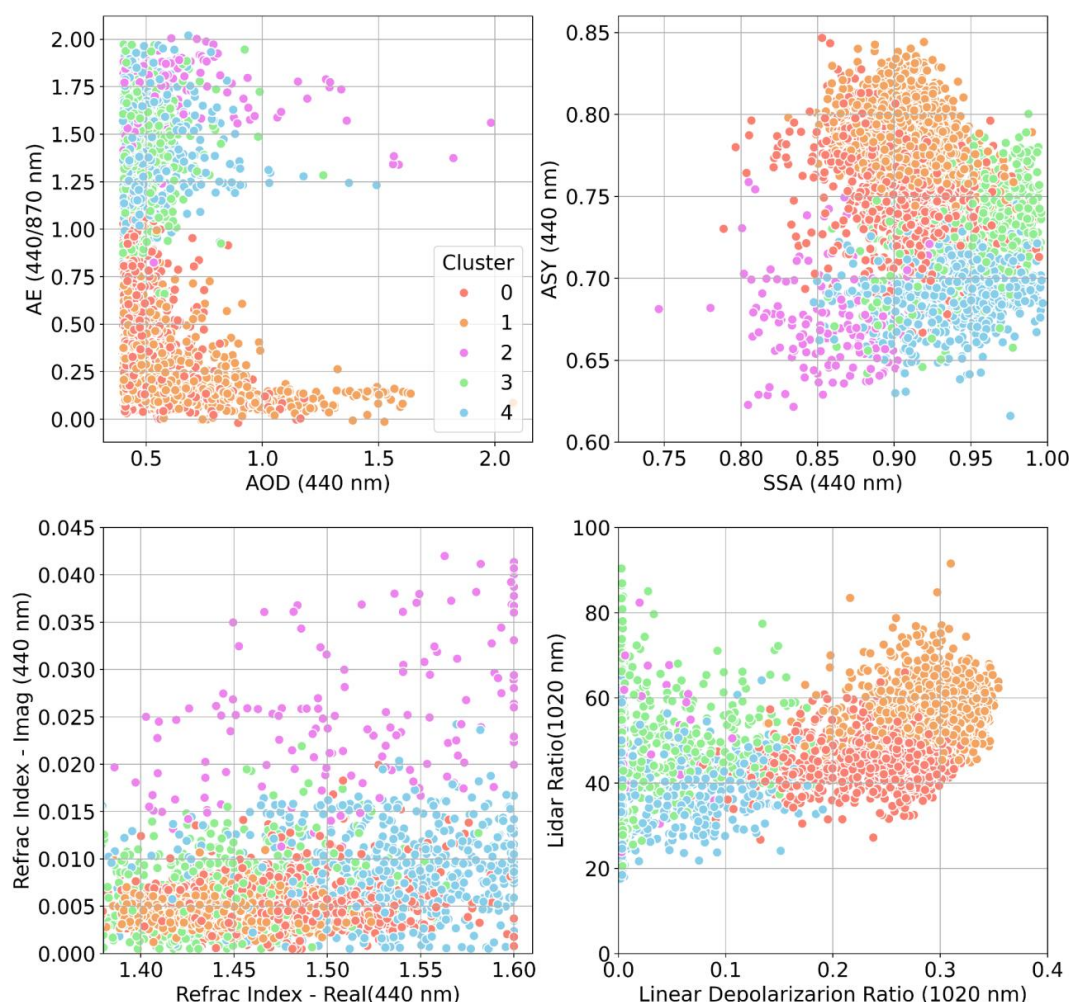
**Figure 2:** Average of sum of squares within-cluster and cluster stability as function of the number of clusters.

We applied the cluster analysis once we defined the optimal number of clusters. **Figure 3** presents a combination of graphics used for aerosol properties analysis, highlighting the obtained clusters' behavior and distinction. The first graphic (Fig. 3a) represents the Aerosol Optical Depth (AOD) as a function of Angstrom Exponent (AE), which allows us to relate aerosol loading variability with aerosol-regimes dominated either by coarse or fine mode (Eck et al., 1999). This analysis shows that two of the clusters (C0 and C1) are regimes dominated by coarse mode particles ( $AE < 1.0$ ), while the remaining three (C2, C3, and C4) are regimes under stronger influence of fine mode particles ( $AE > 1.0$ ). The second plot displays the asymmetric parameter against the single scattering albedo at 440 nm. This plot



aims to elucidate the clusters distinctions related to particles absorption efficiency and the asymmetry between hemispherical forward and backward scattering. Aerosol regimes dominated by coarse particles tend to exhibit more significant forward scattering and, consequently, higher asymmetry parameter values. In contrast, lower asymmetry parameter values are expected in fine mode regimes (Eck et al., 1999, Dubovik et al., 2002). This pattern is evident in the graphic; clusters C0 and C1 present higher asymmetry parameter values, It is also possible to identify the distinction between the non-dust regimes C2, C3 and C4. C2 presents the lowest asymmetry parameter values while it is the most absorbing of the clusters, according to its single scattering albedo values. Small and highly absorbing particles are commonly associated with urban pollution or fresh smoke plumes from biomass burning (Dubovik et al., Omar et al., 2005, Levy et al. 2010, Martins et al. 2009). The C3 cluster differs significantly from C2 by presenting higher asymmetry parameter values, an indication of a shift to larger particle sizes. C3 has higher single-scattering albedo values, indicating a less absorbing aerosol regime. SSA alone did not help to differentiate the two clusters dominated by coarse mode particles (C0 and C1). C0 asymmetry parameter values tend to be lower than those of C1, suggesting that the former could be a dusty mixture not as close to a pure dust scenario as C1. The traditional plot of Lidar Ratio (LR) against Linear Depolarization Ratio (LDR) (Kanitz et al. 2013, Illingworth et al., 2015) confirms this hypothesis (Fig. 3d). Pure dust regimes of aerosol, due to its high level of non-spherical particles, produce higher LDR (Groß et al., 2011). The C1 cluster presents higher values of LDR than C0, indicating that C1 is closer to a pure dust regime. The C0, while a dust regime, is likely to represent a mixed scenario given its LDR values consistent with dust and smoke mixing (Kanitz et al. 2013). LDR values below 15%, which is the case of the clusters C2, C3 and C4, are typically associated with fresh/aged smoke, urban-industrial pollution, and marine particles scenarios. The analysis of the real part versus the imaginary part of the complex refractive index (Fig. 3c) emphasizes C2 as the aerosol regime with the largest absorption and highlights that the real part of the complex refractive index is the main aspect differentiating C3 and C4.



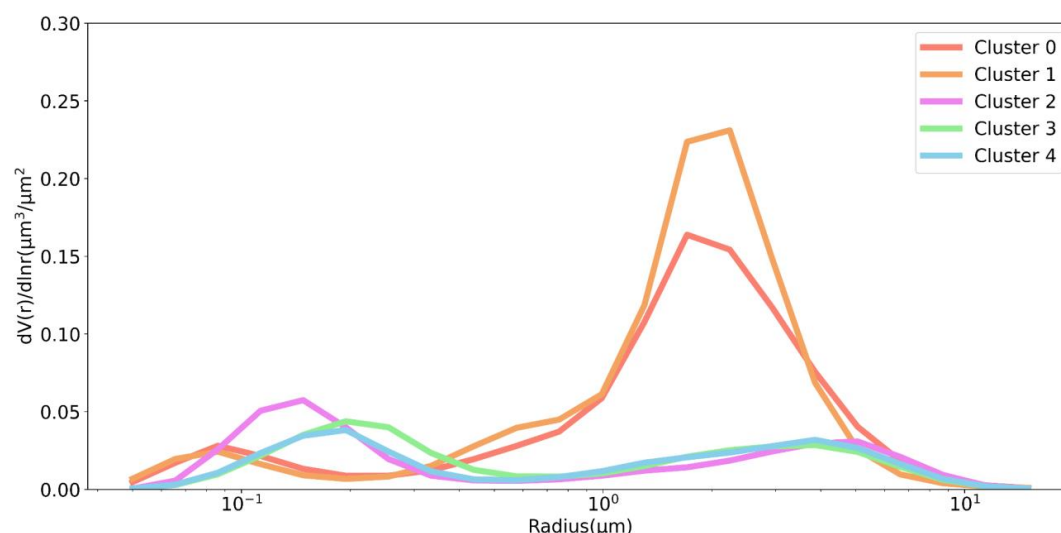


**Figure 3:** Scatterplot of the clusters elements as function of different parameters: a) Extinction Angstrom Exponent (AE) as function of Aerosol Optical Depth (AOD) at 440 nm; b) Asymmetry Parameter (ASY) as function of Single Scattering Albedo (SSA) at 440 nm; c) Lidar Ratio as a function of Linear Depolarization Ratio at 1020 nm; d) Refractive index at 440 nm: Imaginary part as function of Real part.

**Figure 4** and **5** present the clusters average for selected features: size distribution, complex refractive index, single scattering albedo, and asymmetry parameter. A more detailed summary of the mean behavior of the clusters is presented in **Table 3**. The average size distribution of the clusters confirms that aerosol regimes affecting the Iberian Peninsula vary between two scenarios dominated by coarse mode (C0, C1), named here as dust regimes, and three scenarios when coarse mode is not dominant, here considered as non-dust regimes. There are differences between the dust regimes: C1 is associated with a higher coarse particle



loading than C0. Among the non-dust regimes (C2, C3 and C4), the main difference is seen between C2 and the two others. C2 is characterized by larger fine particles loading. Between C3 and C4, one can observe a larger radius spread for C3 regarding the contribution of the fine mode.



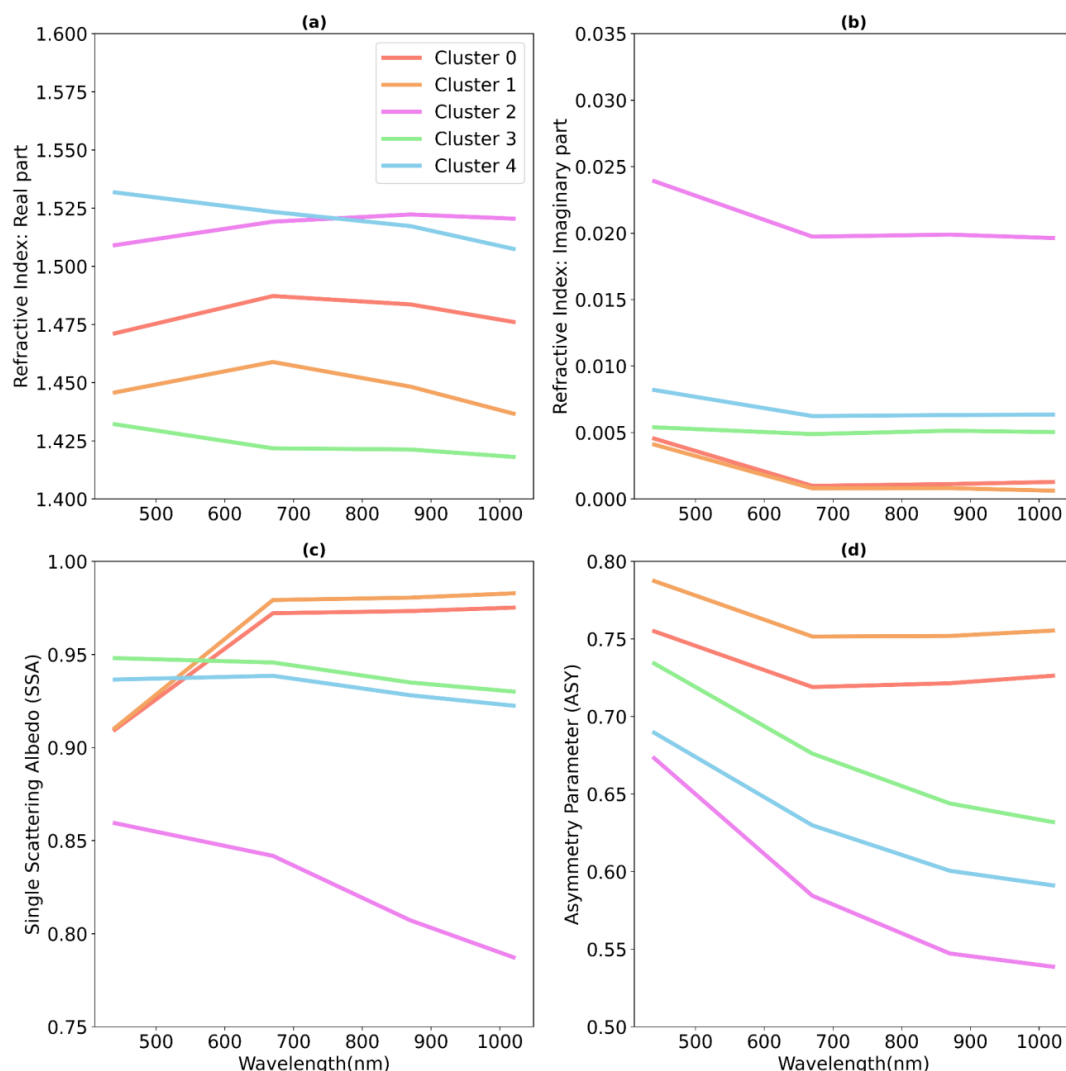
**Figure 4:** Clusters mean volume particle size distribution as a function of radius. The numeric values of each cluster size distribution can be found in Table S2 in the supplement.

Clusters C2 and C4 have close values for the real part of the refractive index but the former C2 has a much larger imaginary part, justifying its lowest SSA (**Figure 5**). The C2 strong absorption combined with its smaller particles suggest that it is likely associated with fresh smoke (Reid et al., 1998, Reid et al., 2005). The average of the real part of the complex refractive index corroborates the difference between the C3 and C4 aerosol regimes. According to Moise et al. (2015), a variation as such observed between C3 and C4 (1.4 to 1.5) could produce an increment of 12 % in estimating the direct aerosol radiative forcing over the solar spectrum wavelength range. Zhao et al. (2019) also showed that the direct aerosol radiative forcing is estimated to vary by 40 % when the real part of the complex index values varies between 1.36 and 1.56. The reasons for the differences observed between the real parts of C3 and C4 remain unclear. However, the spatial distribution of the clusters (see Fig. 6) indicates that C3 is more prevalent in the eastern region of the Iberian Peninsula, which is the wettest area and more exposed to air masses from the Mediterranean and Eastern Europe. Additionally, the low values of the real part of the complex refractive index for C3 align with aerosol regimes that have a strong contribution from sulfate particles. The spectral dependency of the single scattering albedo corroborates our attribution of the C0 and C1 to a dust regime. Dust particles are characterized by strong absorption in the UV spectrum (Dubovik et al., 2002), which decreases as the wavelength increases, a feature present in both



485 C0 and C1. Also, consistent with dust-dominated regimes, C0 and C1 have the largest mean  
486 asymmetry parameter at all wavelengths.

487



488

489 **Figure 5:** Clusters average of complex refractive index, (a) Real and (b) Imaginary parts, (c)  
490 single scattering albedo and (d) asymmetry parameter.

491

492 The analysis above and the summary provided by **Table 3** provide several specific  
493 characteristics that help us to contextualize the clusters. To enhance this understanding, we  
494 add the spatial (**Figure 6**) and seasonal (**Figure 7**) distribution of the clusters into our  
495 analysis. C0 and C1 aerosols regimes are dominated by dust, where C1 is the closest regime



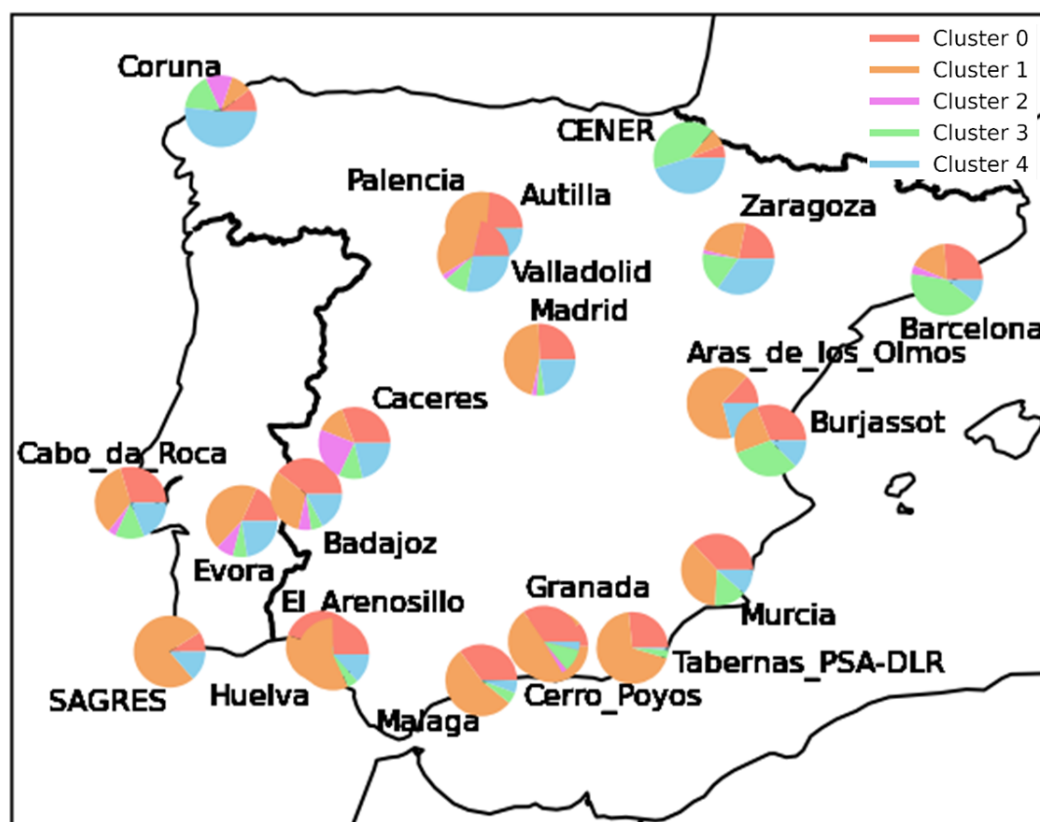
to what we could call pure dust scenario. Both aerosol regimes, C0 and C1, affect practically the entire Peninsula (**Figure 6**) and all year round, but it is more frequent in the southern part of the Peninsula, an expected feature considering that the dust particles are mainly transported from North Africa (Cachorro et al., 2016, Gómez-Amo et al., 2017). The C2 cluster is the most absorbing regime, and is characterized by the smallest fine mode particles. Our hypothesis is that C2 is associated with fresh smoke. Its spatial distribution (**Figure 6**) with more frequent occurrence along the belt spanning from Evora, in Portugal, to Cáceres, in Spain, a region known for high recurrence of biomass burning, reinforces our hypothesis. Additionally, the seasonal distribution of C2 in this region coincides with the peak of the biomass burning season. C3 aerosol regimes also occur over all AERONET sites during all seasons, but it is dominant in the eastern and northeastern portions of the Iberian Peninsula. Among non-dust regimes, its unique feature is its very low real part of the refractive index. C4, as C3, is weakly absorbing according to their single scattering albedo. However, it is present across the entire Peninsula, but its occurrence increases in the central and in the northern portions, which are more prone to biomass burning. An important feature of C4 is that its occurrence increases during the summer and beginning of autumn (**Figure 7**) in the central region of the Iberian Peninsula, from Évora (Portugal) to Madrid (Spain), when the region's biomass burning season is going on. These aspects led us to hypothesize that C4 is an aerosol regime under strong influence of smoke aerosol particles.

**Table 3:** Summary of the clusters based on the average of optical and microphysical properties. A detailed description of the clusters can be found in Tables S1 and S2 in the supplement.

Properties	Cluster0 (Polluted dust)	Cluster1 (Pure dust)	Cluster2 (Strongly absorbing smoke)	Cluster3 (Urban-Industrial Pollution)	Cluster4 (Moderately absorbing smoke)
Number of records	1308	1665	153	660	604
Percentage (%)	29.76	37.88	3.48	15.01	13.74
Ref_Idx_Real (440 nm)	1.47(0.04)	1.44(0.03)	1.51(0.07)	1.43(0.06)	1.52(0.05)
Ref_Idx_Img (440 nm)	0.005(0.002)	0.004(0.001)	0.025(0.009)	0.006(0.004)	0.009(0.004)
VMR-F	0.14(0.03)	0.14(0.03)	0.16(0.02)	0.21(0.04)	0.18(0.04)
STD - F	0.61(0.09)	0.67(0.07)	0.42(0.06)	0.47(0.06)	0.41(0.05)
REff-F	0.12(0.02)	0.12(0.02)	0.14(0.02)	0.18(0.03)	0.17(0.03)
REff-C	1.68(0.16)	1.61(0.13)	2.44(0.43)	2.31(0.38)	2.25(0.49)
VMR-C	2.02(0.23)	1.88(0.17)	3.10(0.45)	2.82(0.42)	2.82(0.57)
STD-C	0.60(0.52)	0.54(0.04)	0.68(0.06)	0.63(0.05)	0.67(0.05)
AOD (440 nm)	0.50(0.11)	0.58(0.21)	0.64(0.29)	0.48(0.09)	0.51(0.13)
SSA (440 nm)	0.91(0.03)	0.91(0.02)	0.86(0.03)	0.95(0.03)	0.94(0.03)
ASY (440 nm)	0.76(0.02)	0.79(0.19)	0.67(0.03)	0.73(0.03)	0.69(0.02)
AE(440/870 nm)	0.40(0.25)	0.24(0.14)	1.67(0.20)	1.43(0.26)	1.47(0.25)
LR(1020 nm)	64(9)	70(8)	89(16)	77(17)	61(15)
LDPR(440 nm)	0.17(0.04)	0.21(0.04)	0.01(0.03)	0.03(0.04)	0.03(0.05)



520



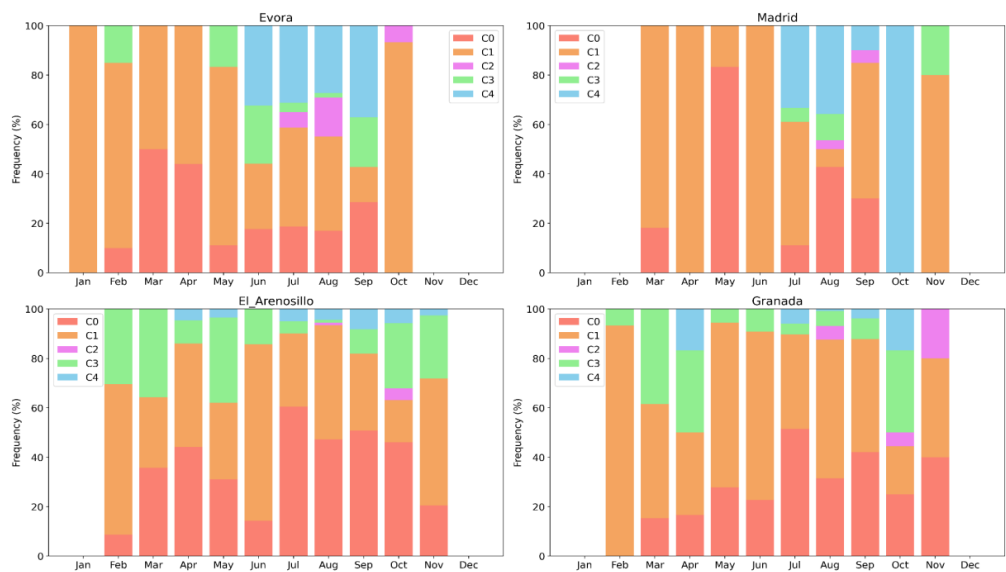
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522 **Figure 6: Proportions** of the occurrence of the clusters of aerosol regimes at the AERONET  
523 sites across the Iberian Peninsula.

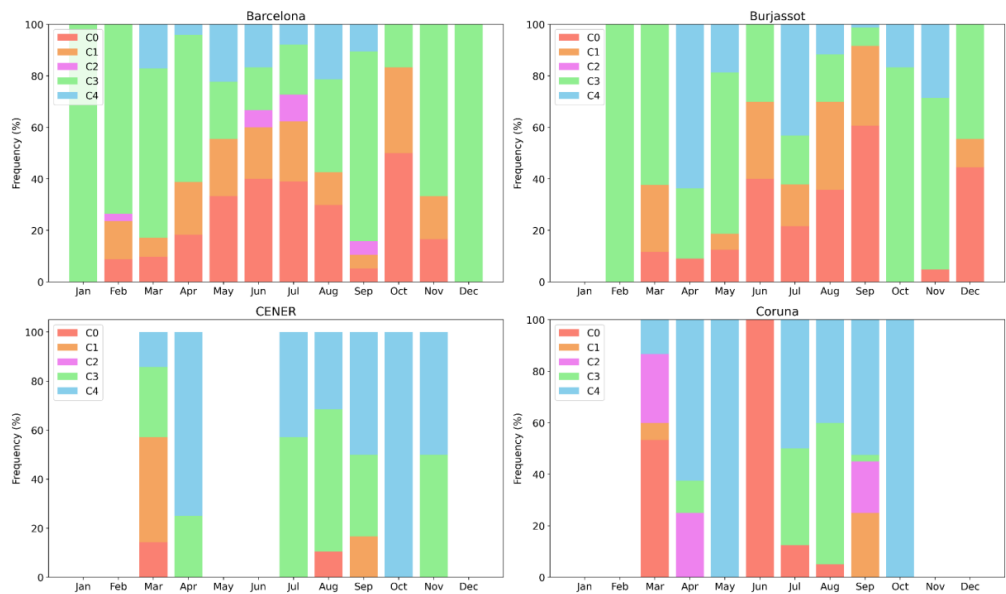
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525 **Figures 7 and 8** provide a perspective view on the seasonal occurrence of each cluster based  
526 on sites that represent different regions of the Iberian Peninsula.





**Figure 7:** Clusters relative monthly occurrence over the AERONET sites representatives of the Iberian Peninsula western lowlands (Evora), highlands plateau (Madrid) and southeast lowlands (El Arenosillo, Granada).



**Figure 8:** Clusters relative monthly occurrence over the AERONET sites representatives of the following Iberian Peninsula regions: Eastern Coast (Barcelona, Burjassot) and Northern (Coruna, CENER).

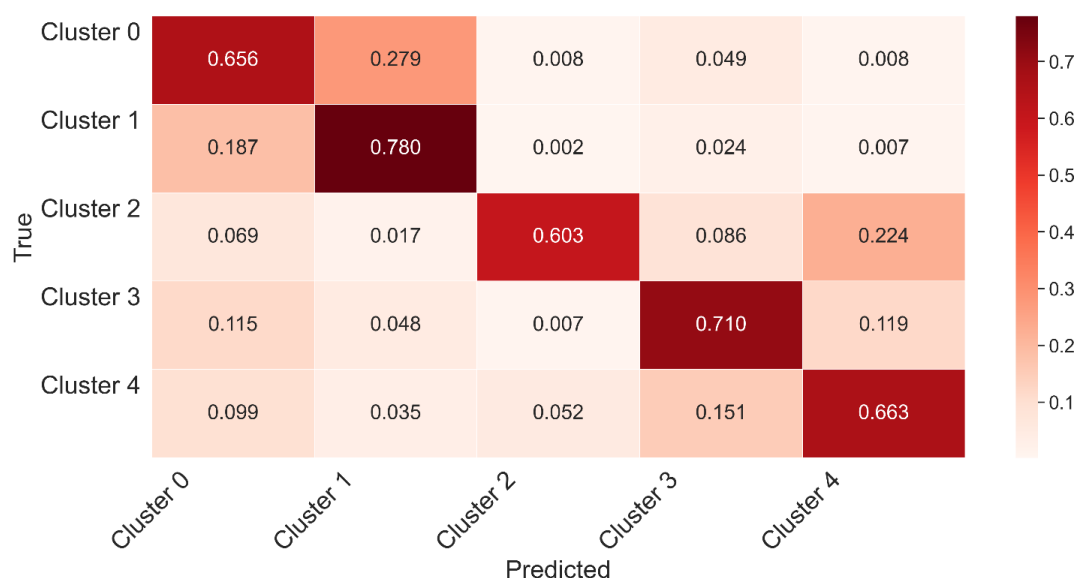


### 535 3.2 Random Forest Classifier: Performance and Optical models spatial dynamic

536 The Random Forest training of MERRA-2 aerosol-type column mass density as predictors of  
 537 aerosol optical regime covered 70% of the AERONET sky inversions used in this study, combining  
 538 datasets from all sites. The testing dataset, constituted by the remaining 30%, was used to evaluate  
 539 the model performance. The best parameters obtained from the optimization using  
 540 RandomizedSearchCV were the number of decision trees of 477 ( $n\_estimators = 477$ ) and  
 541 maximum depth of trees of 19 ( $max\_depth=19$ ). There are several metrics for accessing machine  
 542 learning performance. **Figure 9** presents the one used in this study, the Normalized Confusion  
 543 Matrix (NCM), which expresses the percentage of correct and incorrect predictions (where  
 544 the classifier got confused). In the matrix, the rows represent the true labels, and the columns  
 545 represent the predicted ones. The values along the diagonal indicate the percentage of times  
 546 where the predicted matches the true label. The other cells reflect instances where the  
 547 classifier mislabeled an observation; the column tells us what the classifier predicted, and the  
 548 row tells us the correct label.

549 For all clusters, the classifier's correct predictions surpassed the incorrect predictions, with  
 550 a maximum frequency of correct prediction close to 80% obtained for C1. The minimum  
 551 percentage of correct prediction, about 60%, was obtained for C2, the highest absorbing  
 552 cluster. Regarding dust regime clusters, despite the struggle to predict C0, it is possible to see  
 553 that, in this case, the classifier's main confusion is with the C1, which is also a cluster related  
 554 to an aerosol scenario dominated by coarse mode particles (dust regime), as C0. Therefore,  
 555 this is a somehow expected confusion, which would not introduce a substantial error in the  
 556 radiative effect calculations. Rarely does the classifier take either C0 or C1 as C2, C3, and C4,  
 557 a case where substantial error in the radiative effect would be expected. By combining C0  
 558 and C1 results in the NCM, the percentage of correct prediction achieved by the classifier  
 559 indicating dust regime is higher than 95%. Similarly, rarely the classifier takes C3 and C4 as  
 560 C0, C1 and C2. Given that C3 and C4 are also close in terms of their optical properties,  
 561 especially concerning absorption, some degree of confusion among them is  
 562 expected. Nevertheless, these aspects of the confusion matrix among close clusters are  
 563 important to identify where the model needs extra training. C2, the less frequent and the one  
 564 representing the most absorbing aerosol regime over the Iberian Peninsula is rarely  
 565 mislabeled as C0 or C1, but often mislabeled as C3 or C4. Still, the score percentage is around  
 566 60%.

567



568

569 **Figure 9:** Normalized confusion matrix of the Random Forest classifier applied to the prediction  
570 of the clusters that describe the typical aerosol optical regime based on MERRA-2 aerosol  
571 components column mass density.

572 To provide further insight into the model performance, we also examined other metrics  
573 commonly used to evaluate Random Forest training: Accuracy, Precision, Recall, and F1 score  
574 (Table 4). The trained model achieved a general accuracy of 70 %, meaning it correctly  
575 predicted the aerosol regime in three out of four cases. For all clusters, all metrics adopted  
576 were higher than 0.60, with precision and recall values exceeding 0.75 in some cases. The  
577 precision metric indicates how often the positive predictions are correct. The model  
578 precision varied within the specific optical regimes (ex., non-dust) and among optical  
579 regimes (dust, non-dust). It showed higher precision in identifying C1 than C0, the two dust-  
580 regimes. Among the non-dust regime clusters, the highest precision obtained was related to  
581 the prediction of C2, suggesting a lower likelihood of false positive for this class given its  
582 strong absorption nature, mislabeling this aerosol regime would translate in high cost due to  
583 significant radiative error; therefore, its highest precision is a promising outcome.

584

585 **Table 4.** Performance metrics values of the trained model prediction of aerosol optical  
586 regime based on aerosol-type column mass density.

Clusters	Precision	Recall	F1-Score	Support (N)
0	0.62	0.62	0.62	394
1	0.68	0.70	0.69	452

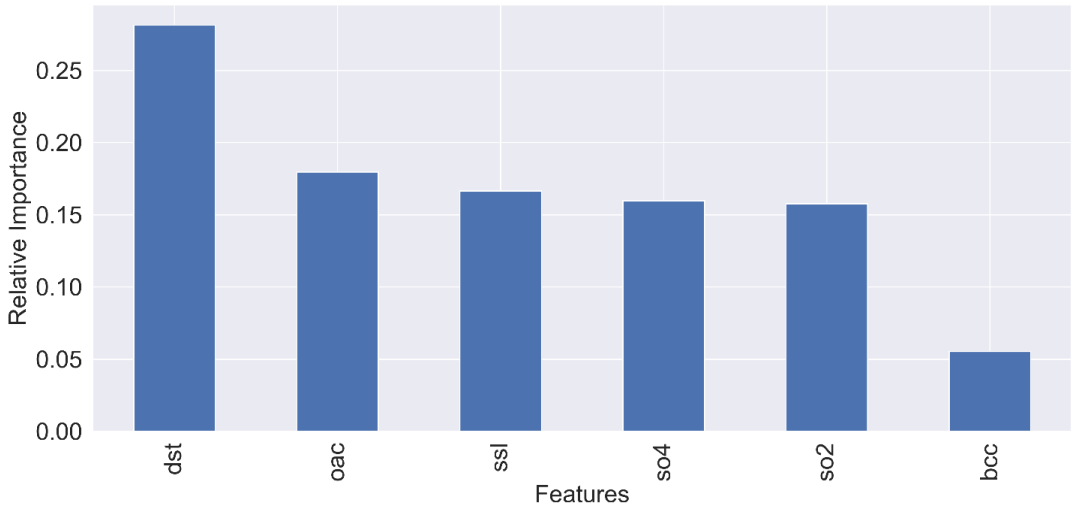


2	0.62	0.60	0.61	62
3	0.76	0.73	0.74	251
4	0.68	0.69	0.69	185

587

588 **Figure 10** illustrates the relative importance of the predictor variables, highlighting the  
589 influence of each aerosol-type column mass density on the model's decision-making. The  
590 results indicate that the presence of dust and organic carbon over the Iberian Peninsula is  
591 the primary factor affecting the aerosol optical properties in this region. This finding aligns  
592 with actual conditions, as the transport of Saharan dust to the peninsula and biomass burning  
593 are the two main sources driving the variability of aerosol optical properties in the area.  
594 Interestingly, black carbon column mass density did not rank among the top predictors.  
595 Despite the expectation that black carbon might serve as a significant indicator of the aerosol  
596 optical regime due to its association with smoke-influenced aerosols, there is considerable  
597 uncertainty in black carbon simulations in atmospheric chemistry models, including  
598 reanalyses such as MERRA-2. This uncertainty may hinder its effectiveness in predicting the  
599 aerosol regime observed at AERONET monitoring sites.

600



601

602 **Figure 10:** Relative importance of the predictor variables, i. e. the degree of influence of each  
603 aerosol-type column mass density on the model decision-making.

604



### 3.3 Application: Case studies

From the testing dataset, we selected some case studies that significantly impacted local populations, garnered media attention, and represented different aerosol scenarios in the Iberian Peninsula. This selection provides a visual (qualitative) demonstration of the model's predicting capability (**Table 5**).

**Table 5:** List of case studies of aerosols high loading events over Iberian Peninsula selected to highlight as examples of the classifier trained model application.

Case study	Date	Nature (Reference link)
#01	June 27, 2023	Smoke <sup>1</sup>
#02	October 16, 2017	Dust and Smoke <sup>2</sup>
#03	August 11, 2016	Smoke <sup>3</sup>
#04	March 17, 2022	Dust <sup>4</sup>

1-<https://earthobservatory.nasa.gov/images/151507/canadian-smoke-reaches-europe>

2-<https://atmosphere.copernicus.eu/saharan-dust-and-smoke-over-france-and-uk>

3-<https://earthobservatory.nasa.gov/images/88552/fires-rage-in-portugal>

4- <https://earthobservatory.nasa.gov/images/149588/an-atmospheric-river-of-dust>

We set our trained model to prescribe the spatial distribution of aerosol optical regimes (clusters) that best fit various scenarios based on MERRA-2 aerosol-type column mass density. The results for all cases studied are presented in **Figure 11**. Since AERONET sky inversion products only provide a complete characterization of aerosol microphysical (size distribution plus complex refractive index) and optical properties (Asymmetry parameter and Single Scattering Albedo) for conditions of AOD at 440 nm exceeding 0.4, we will only discuss the optical regime prescriptions for areas where AOD was above this threshold.

For our analysis, we used the MERRA-2 AOD field as a reference.

Case#01 occurred from June 1 to 25, 2023, coinciding with large-scale wildfire events in Quebec, Canada. A substantial portion of smoke from these wildfires crossed the Atlantic Ocean and reached Western Europe, especially the Iberian Peninsula, resulting in darkened skies in the affected countries. Our trained model predicted that the most suitable aerosol optical regime for the areas impacted by the smoke (Portugal, Western, and Northern Spain) is C4, which corroborates our previous discussion associating the C4 optical regime to regional smoke.

Case#02 features an emblematic event on October 16, 2017, marked by simultaneous massive wildfire in central and northern Portugal and a strong dust transport from North

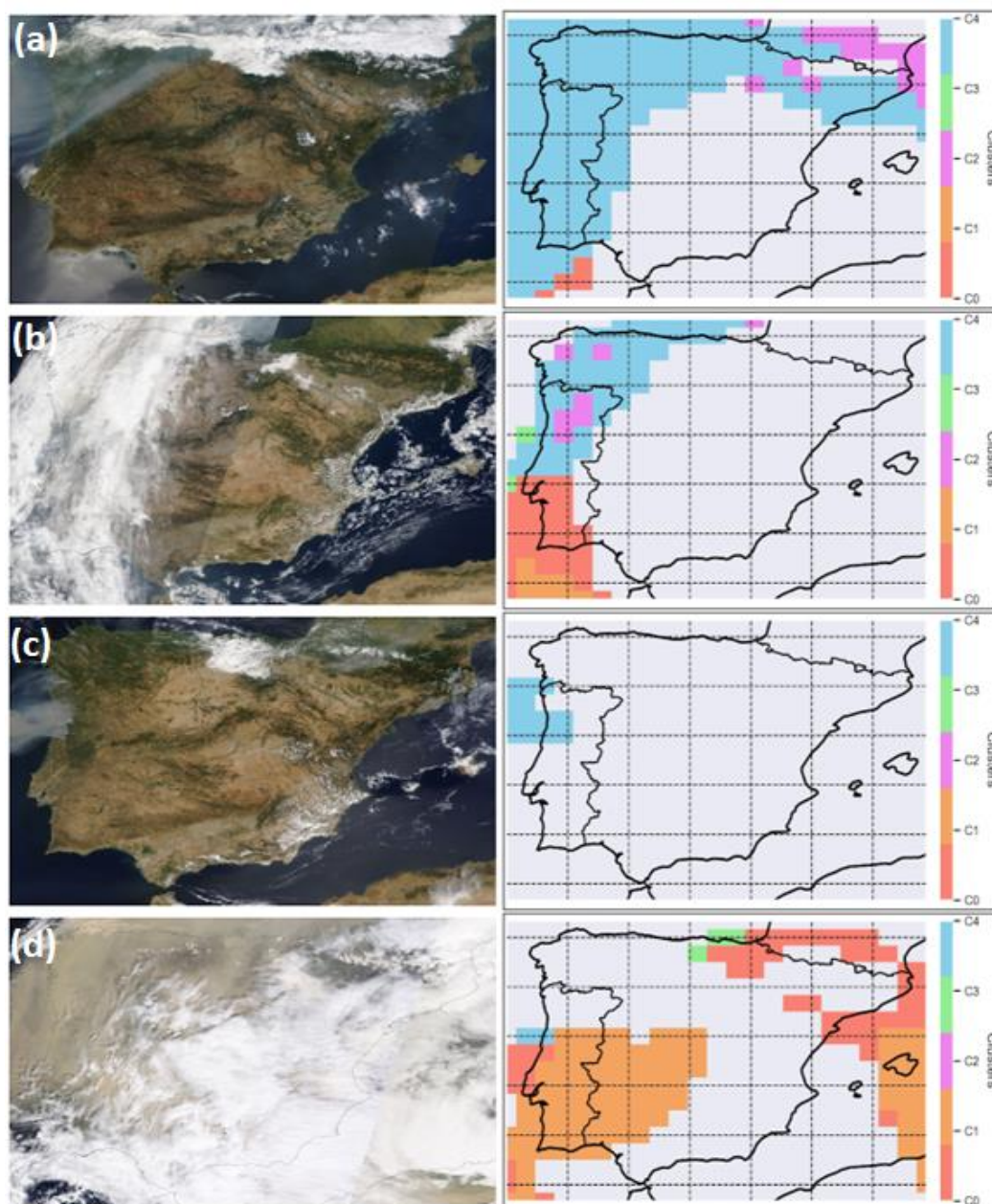




635 Africa via the south of Portugal. The corridor connecting the smoke and dust produced a  
636 strong northward transport affecting the United Kingdom, influenced by the synoptic  
637 conditions associated with the ex-hurricane Ophelia, located just north of the Iberian  
638 Peninsula (Osborne et al., 2019). The optical regime prescription identified the C4 cluster as  
639 the appropriate regime from central Portugal northward to the UK. Meanwhile, the area  
640 affected by dust, spanning from North of Africa to southern and central Portugal, was  
641 characterized by a mix of C0 and C1, the clusters associated with dust regimes. As the dust  
642 plume arrived in Portugal, the model indicated a gradual transition from C1, indicative of  
643 pure dust, to C0, which represents conditions of dust mixed with smoke (Gómez-Amo et al.,  
644 2017). The random distribution of C2 within the larger C4 regions likely reflects the model's  
645 response to the specific conditions dictated by the aerosol-type column mass densities. This  
646 could suggest patches of high-absorbing aerosol-type within a less-absorbing large-scale  
647 smoke plume, although there is insufficient evidence to draw definitive conclusions.

648 Case#03, dated August 16, 2016, involved strong wildfire emissions in northern Portugal.  
649 Most of the smoke was transported toward the Atlantic Ocean, while the remainder of the  
650 peninsula experienced low aerosol loading conditions. Consistent with smoke aerosol  
651 scenarios, the model prescribed the C4 optical regime.

652 Case#04 pertains to an extreme Saharan dust transport that affected most of the Iberian  
653 Peninsula on March 15-17, 2022. During this event, the 24-hour average concentration of  
654 PM<sub>2.5</sub> reached as high as 700  $\mu\text{g m}^{-3}$  in parts of Spain (Rodríguez and López-Darias, 2024).  
655 The pollution episode was dominated by dust, and indeed, the model prescribed the optical  
656 regime C1, which indicates the pure dust conditions for most of the Iberian Peninsula. This  
657 demonstrates our approach's capability to differentiate specific scenarios within dust  
658 regimes. For non-dust regimes regarding C2, a highly absorbing regime, we would not expect  
659 to see widespread prescriptions, as we hypothesize that it is associated with fresh, high-  
660 absorbing pollution plumes. **Figure 6**, depicting the occurrence of each cluster across the  
661 Iberian Peninsula, corroborates our hypothesis by indicating that the C2 regime is mainly  
662 present in specific areas where aerosol loading increases are primarily attributed to biomass  
663 burning, such as the western lowlands of Iberian Peninsula (Evora, Badajoz, and Caceres) and  
664 in the Galicia region (Coruna). The C3 optical regime was not linked to large-scale dust  
665 transport or smoke plumes across the Iberian Peninsula, suggesting it might be associated  
666 with high levels of local or regional pollution. **Figure 6** shows that the C3 regime is commonly  
667 observed throughout the year in the eastern portion of the Iberian Peninsula. The results of  
668 these case studies, combined with performance evaluations, highlight the capability and  
669 potential of this machine-learning approach, which uses clustering and random forest  
670 classification to prescribe optical models from aerosol-type columnar mass density to  
671 calculate aerosol particles' direct radiative effect in atmospheric models. By constraining  
672 modelling with observational data, we can help mitigate the known uncertainties related to  
673 aerosol direct radiative forcing. Additionally, our method's straightforwardness and lower  
674 computational cost favor operational modeling when infrastructure is limited.



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**Figure 11:** Case studies of distinct aerosol scenarios over the Iberian Peninsula selected to test our machine-learning based approach to predict the best optical property regime: (a) Case#01 on June 27, 2023; (b) Case#02 on October 16, 2017; (c) Case#03 on August 11, 2016; (d) Case#04 on March 15, 2022. On the left side, MODIS/NASA True color satellite images (<https://wvs.earthdata.nasa.gov>); and on the right the cluster spatial distribution prescribed by the model.



**Figure 12** shows the single scattering albedo at 550 nm field, comparing the current approach and MERRA-2 reanalysis results. The MERRA-2 columnar total SSA was calculated based on the ratio of total scattering aerosol optical depth to total extinction aerosol optical depth. For smoke scenarios on June 27, 2023, MERRA-2 indicated a more absorbing optical regime (SSA at 550 nm  $\sim 0.86 - 0.90$ ) compared to the current approach (SSA at 550 nm  $\sim 0.95$ ). On this day, the average SSA at 550 nm over the AERONET site in Coruna City, which was directly affected by Canadian smoke, exceeded 0.95. A similar trend was observed for the dust scenarios. For example, on March 17, 2022, the current approach prescribed a less absorbing optical regime (SSA at 550 nm  $\sim 0.94 - 0.95$ ) compared to MERRA-2, which reported a SSA at 550 nm of roughly 0.92 - 0.94. The analysis of SSA at 550 nm over AERONET sites affected by the dust event surpassed 0.94. While these cases highlight differences between the prescriptions based on the clusters and MERRA-2 results, they are only sufficient to warrant further investigation. To gain a statistical perspective on whether the findings from these case studies are isolated incidents or indicative of a trend, we compare a much larger sample of MERRA-2 SSA at 550 nm across various AERONET sites in the Iberian Peninsula using the clusters approach. We focused only on MERRA-2 aerosol scenarios for AOD at 550 nm larger than 0.3 and conducted the comparison segmented by the optical regimes defined by the clusters.

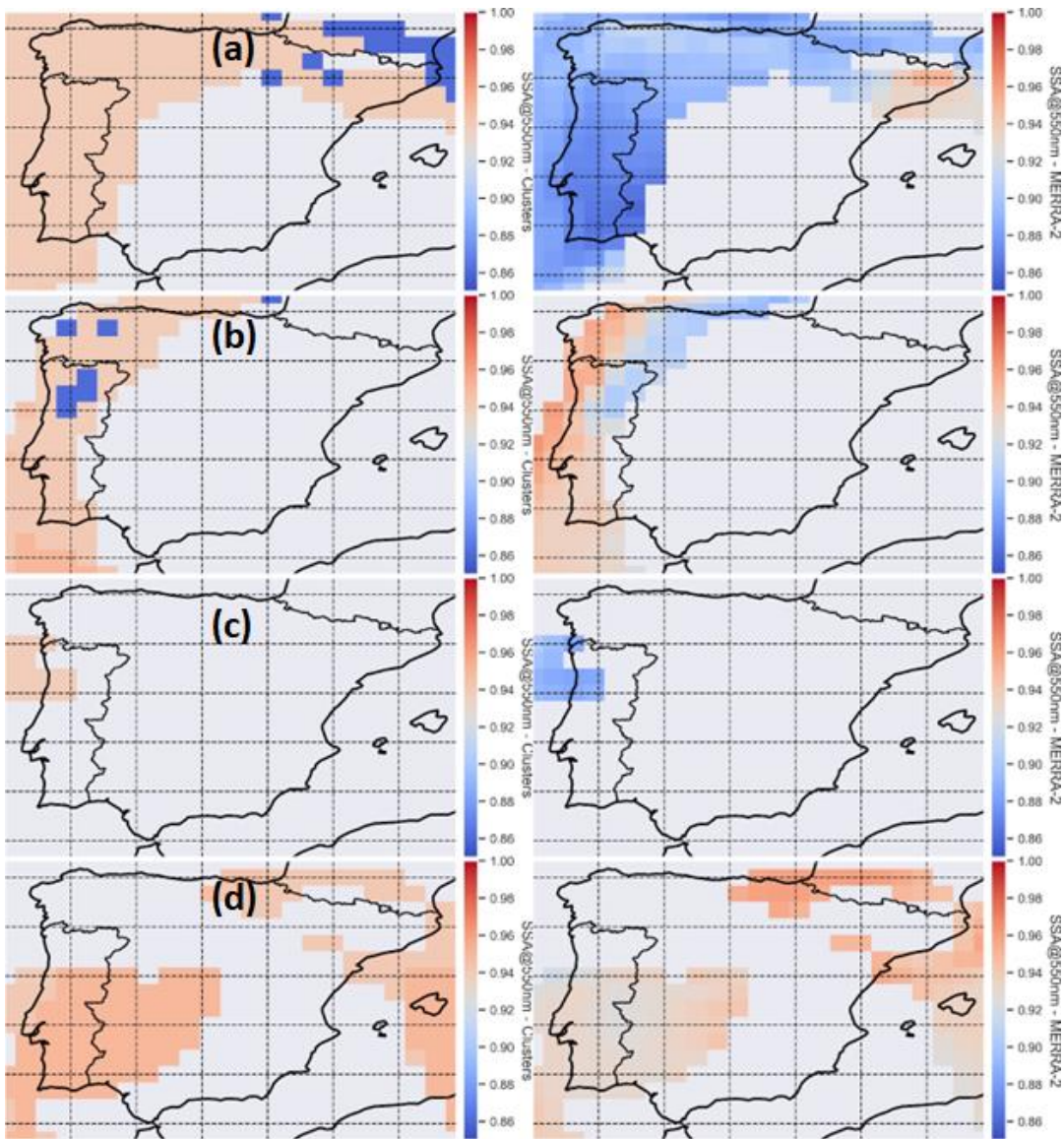
**Figure 13** shows the count distribution of MERRA-2 SSA at 550 nm for the aerosol regimes represented by the clusters C0, C1, C3 and C4, as classified by the random forest classifier we developed. Each cluster's SSA at 550 nm histogram was randomly simulated following a Gaussian distribution considering the cluster mean and standard deviation. A similar analysis was conducted for the Angstrom Exponent (**Figure 14**). In addition to the absorption, we evaluated aspects of mean particle size behaviors. Based on **Figure 13**, we found that, on average, our aerosol optical regime prescription based on the clusters (AERONET) is less absorbing than MERRA-2 for aerosol regimes C0, C1, C3 and C4. More significant differences are observed for C1, C3 and C4. Cluster C1 corresponds to a dust scenario closer to pure dust, while C4 is dominated by smoke. Regarding the particle size indicator (AE), it was observed that MERRA-2 has a lower contribution of coarse particles in the dust regimes compared to the cluster-based prescriptions (**Figure 14 a, b**). This finding aligns with Adebisi et al. (2023), which noted that climate models tend to underestimate large dust particles, mainly when representing North African dust plumes. Conversely, for the non-dust regimes (C3, C4), MERRA-2 shows a larger relative contribution of coarse particles than the clusters-based prescription (**Figure 14 c, d**). **Figure 15** shows the results for C2. For this specific regime, on average, prescription based on the cluster (AERONET) is more absorbing than MERRA-2, opposite to the findings of the other clusters. Regarding AE, under the C2 regime, MERRA-2's mean AE is lower than that prescribed from the cluster, suggesting a lower relative contribution of fine mode in the reanalysis simulations. This is like the findings related to the two other fine mode dominant regimes (C3 and C4).

As demonstrated by the SSA and AE distributions (Figures 13, 14, 15) and evaluated from Table 1 and 4, the model can also predict the occurrence of the minority cluster C2 (3–4 percent of samples). The model preserves the physical distribution characteristics of less frequent aerosol regimes while capturing its features without the need for explicit class imbalance treatment. With C2's highly absorbing and dominant fine mode conditions



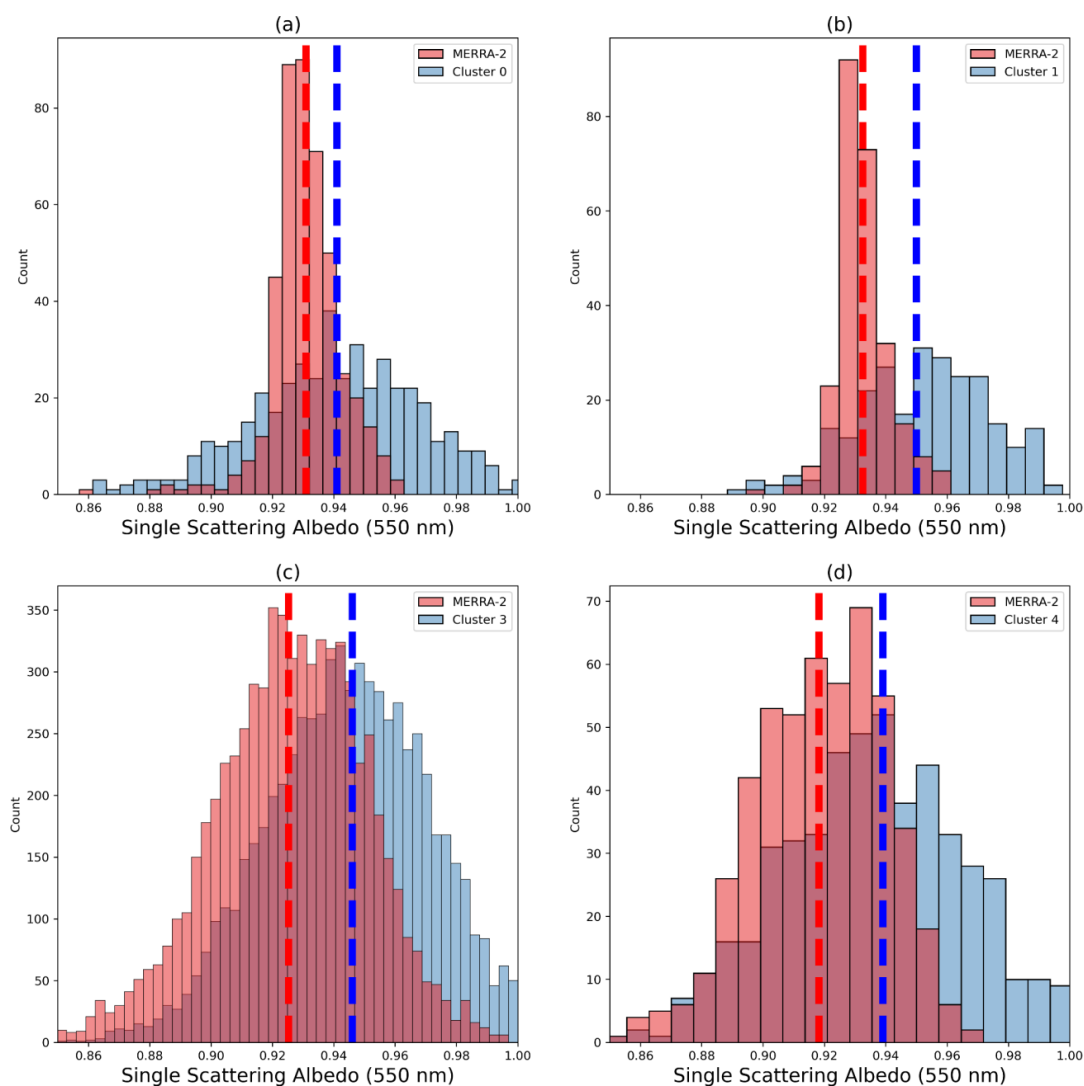


726 reflected in both SSA and AE predictions, the distributions across clusters demonstrate  
727 agreement between expected and observed distribution values.  
728



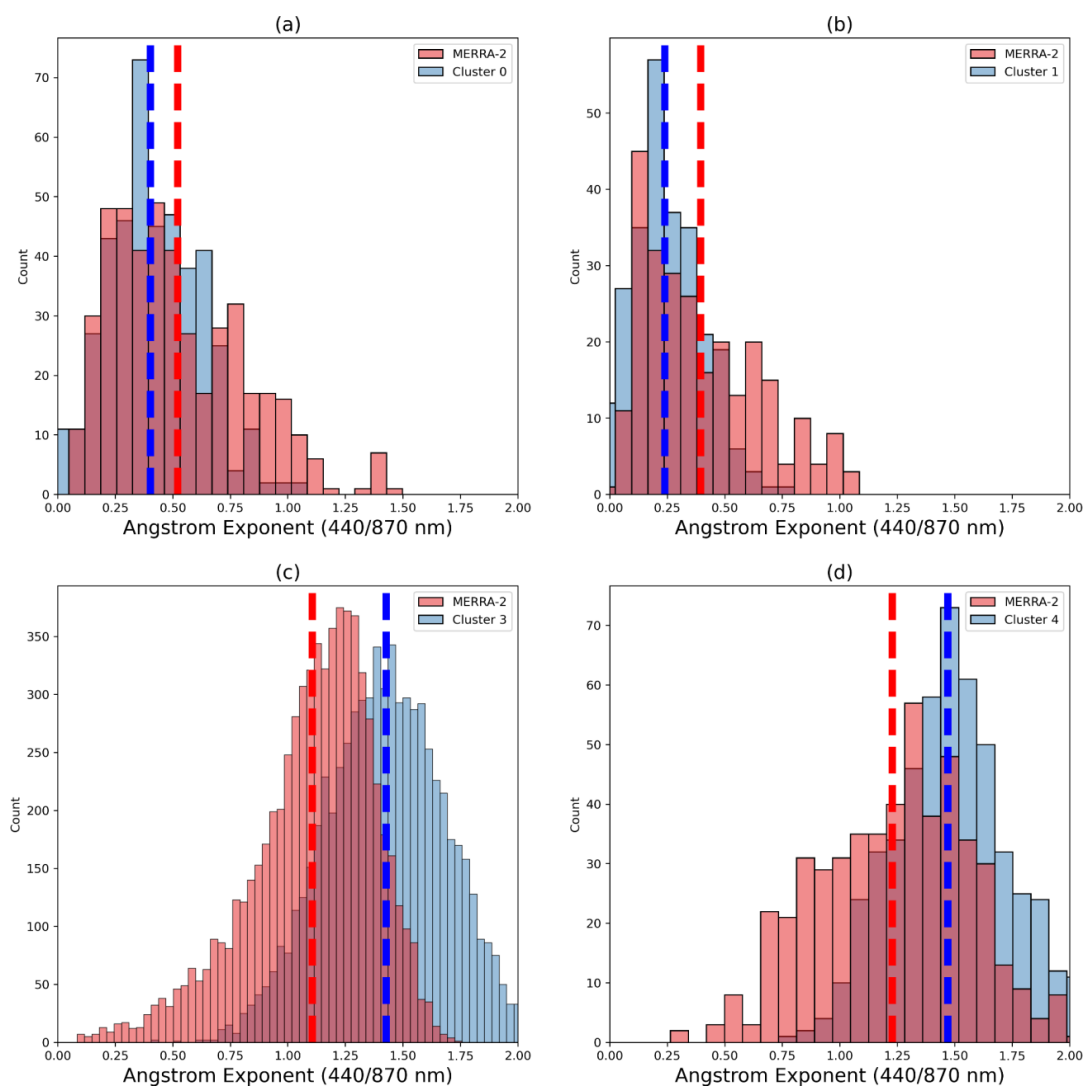
729  
730 **Figure 12:** Single Scattering Albedo (SSA) prescription based on the current study approach  
731 (left) and that simulated by MERRA-2 (right) for the selected case studies of Table 2: (a)  
732 Case#01 on June 27, 2023; (b) Case#02 on October 16, 2017; (c) Case#03 on August 11, 2016;  
733 (d) Case#04 on March 15, 2022.

734

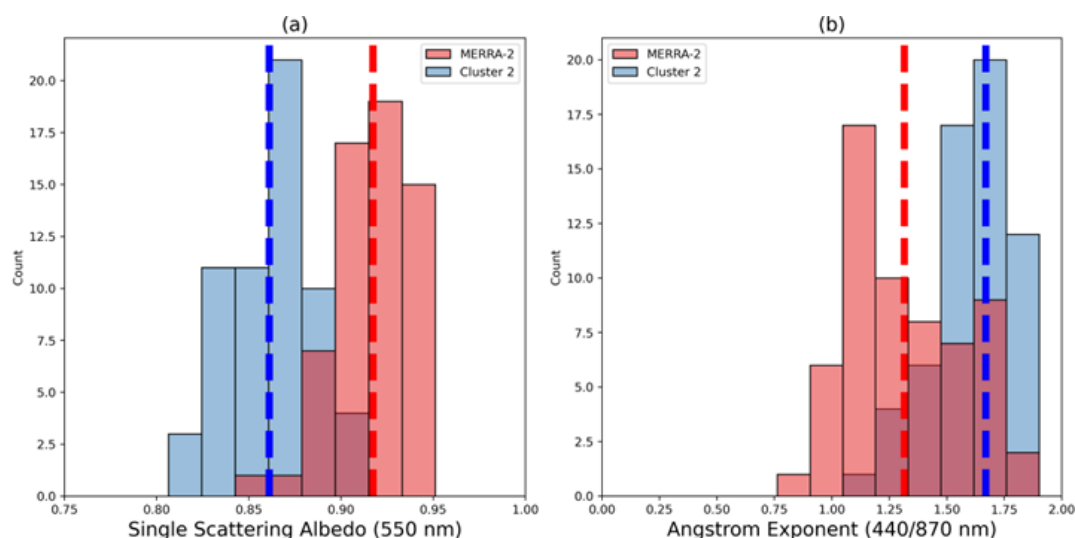


**Figure 13:** Current study prescription and MERRA-2 simulation of Single Scattering Albedo (SSA) frequency distribution as function of the optical regime (clusters): a) Cluster 0; b) Cluster 1; c) Cluster 3; e) Cluster 4.





**Figure 14:** Current study prescription and MERRA-2 simulation of Angstrom Exponent (AE) frequency distribution as function of the optical regime (clusters): a) Cluster 0; b) Cluster 1; c) Cluster 3; e) Cluster 4.



**Figure 15:** Current study prescription and MERRA-2 simulation of (a) Single Scattering Albedo and (b) Angstrom Exponent (AE) frequency distribution for the Cluster 2 scenario.

## 4. Conclusions

This study emphasizes the importance of observational-based research to constrain the prescription of aerosol-intensive properties in atmospheric models. We aimed to characterize the typical aerosol intensive optical properties affecting the Iberian Peninsula (IP) using data from the atmospheric column AERONET sky inversion products. We employed K-means clustering to analyze historical aerosol intensive properties across all AERONET sites that operated for at least two years and had the highest quality dataset level (2.0) available. We identified five distinct clusters (C0, C1, C2, C3, C4) representing different optical regimes, illustrating the predominant aerosol scenarios in the IP. The key difference among these clusters lies in the contribution of coarse mode particles and their absorption efficiency. Clusters C0 and C1 are dominated by coarse mode particles and classified as dust-regimes due to their association with Saharan dust transport. In particular, the optical properties of C1 closely resemble a pure dust scenario, while C0 indicates a more mixed situation, which we refer to as dusty. On the other hand, clusters C2 and C4 are identified as non-dust regimes, linked to strong and moderate absorption related to smoke plumes. Cluster C3, also a non-dust regime, is more frequently observed in the eastern part of the IP and differs from C4 mainly by having a much lower real part of the refractive index. After identifying the typical aerosol regimes affecting the IP, we utilized aerosol-type columnar mass density data (dust, organic carbon, black carbon, sea salt, and sulphates) from MERRA-2 to predict the aerosol optical regime at each grid point using the supervised learning



methodology Random Forest. We tested the performance of the trained model under various aerosol scenarios. The accuracy of the predictions for the aerosol optical regimes ranged from 60% to 75%, depending on the regime, with an average accuracy of 70%. Notably, the accuracy exceeded 90% when predicting solely dust or non-dust optical regimes.

An analysis of MERRA-2 simulations alongside this study's AERONET cluster-based prescriptions of optical regime indicators, such as absorption (SSA) and size (AE), reveals that MERRA-2 is generally more absorbing for the aerosol optical regimes (C0, C1, C3 and C4) impacting the atmosphere of the Iberian Peninsula, except for the most absorbing regime (C2). Specifically, the reanalysis simulations indicate higher absorption under the non-dust regimes C3 and C4. When examining the relative contributions of fine and coarse modes, the cluster-based prescription indicates a larger average contribution of coarse particles than the MERRA-2 under dust-regimes (C0, C1). Conversely, for the non-dust regimes (C2, C3, C4), MERRA-2 shows a lower relative contribution from the fine mode compared to the clusters-based prescription.

Our findings contribute to enhancing the understanding of the dynamic aerosol optical properties over the Iberian Peninsula and highlight the potential of machine-learning approaches to improve the representation of aerosol radiative forcing in atmospheric models. Many atmospheric modelling systems are not designed to simulate aerosol-intensive microphysical and optical properties in real time. Additionally, computational cost remains a common limitation worldwide. Our approach integrates AERONET-derived intensive properties based on climatological optical regimes to refine the model, coupled with predicted aerosol-type columnar mass density. This integration can help reduce regional uncertainty in the simulation of aerosol radiative forcing.

## Competing interests

The authors declare that they have no conflict of interest.

## Acknowledgements and financial support

The authors acknowledge the financial support of FCT—Science and Technology Portuguese Foundation, which funded the project FIRESMOKE (<http://doi.org/10.54499/PTDC/CTA-MET/3392/2020>) through national funds. Thanks are also owed to the financial support given to CESAM by FCT (UID Centro de Estudos do Ambiente e Mar (CESAM) + LA/P/0094/2020) through national funds. We also acknowledge the financial support of CNPq - National Council for Scientific and Technological Development (CNPq) through the funding processes CNPq Nº 441851/2023-1 and CNPq Nº 172486/2023-8. Author HFCV also thanks to the CNPq grant No 315349/2023-9. We thank AERONET and MERRA-2 PIs and teams for their effort in establishing and maintaining the sites and the reanalysis development used in this study. We acknowledge the use of imagery from the Worldview Snapshots application (<https://wvs.earthdata.nasa.gov>), part of the Earth Observing System Data and Information System (EOSDIS).



## 808 Author contributions

809 NR, KL and PT designed and performed the research, analyzed the data, and wrote the first  
810 version of the paper. MY, SF, LF, OM, HFCV contributed to writing, discussion, review and  
811 editing. ICM and AIM conceptualization and coordination of the Project FIRESMOKE,  
812 discussion, review and editing.

## 813 Code and data availability.

814 All the datasets (AERONET and MERRA-2) used in this study are publicly available and were  
815 downloaded from their respective websites (<https://aeronet.gsfc.nasa.gov/>; and  
816 <https://disc.gsfc.nasa.gov/datasets?project=MERRA-2>). Code and dataset required to  
817 conduct the analyses herein is available at <https://doi.org/10.5281/zenodo.15178347>  
818 (Rosario, 2025).

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