

Response to the comments from Referee-1

We sincerely appreciate your thoughtful questions and valuable suggestions, especially the comments regarding wave reflection events during weak polar vortex states or SSWs. We have carefully addressed each of your comments point by point (highlighted in blue) and made the corresponding revisions in the revised version of the manuscript.

Note that there are five main modifications in the revised manuscript: 1) a new structure of the introduction section; 2) a new Table 1 summarizing the analyzed experiments in this study; 3) a new Figure 2 showing the period in days from 6 to 12; 4) a new Figure 6 showing the longitude-pressure cross section of Plumb flux (F_x , F_z); 5) a new Figure 10 showing the zonal wind anomalies at 321 hPa.

General comments

My main concern is that the authors might conflate potentially related yet distinct concepts. The manuscript claims that the results indicate the surface response to stratospheric wave reflection events is causally forced by stratospheric perturbations (L17). Planetary wave reflection is a generally well-defined phenomenon in which planetary waves propagate upward from the troposphere into the stratosphere and are reflected back. Intuitively, upward wave propagation, a reflective surface, and downward wave propagation would be expected in a wave reflection event. The present experiment only shows the downward wave propagation, yet it seems to implicitly imply that downward wave propagation equals wave reflection. I think a discussion is warranted to discuss the model's limitations and clarify the relationship between the results and wave reflection. The authors may also consider softening some related arguments regarding the implications of the results.

We agree that our methodology is not intended to study the entire lifecycle of wave reflection events - indeed the entire point is to isolate the downward propagation part from the antecedent upward pulse, as the antecedent upward pulse necessarily entails strong tropospheric anomalies. In general, we tried to be careful in our wording, but L17 needed to be clarified. Sorry for the confusion. But we want to make clear that there is unambiguous wave reflection in these experiments.

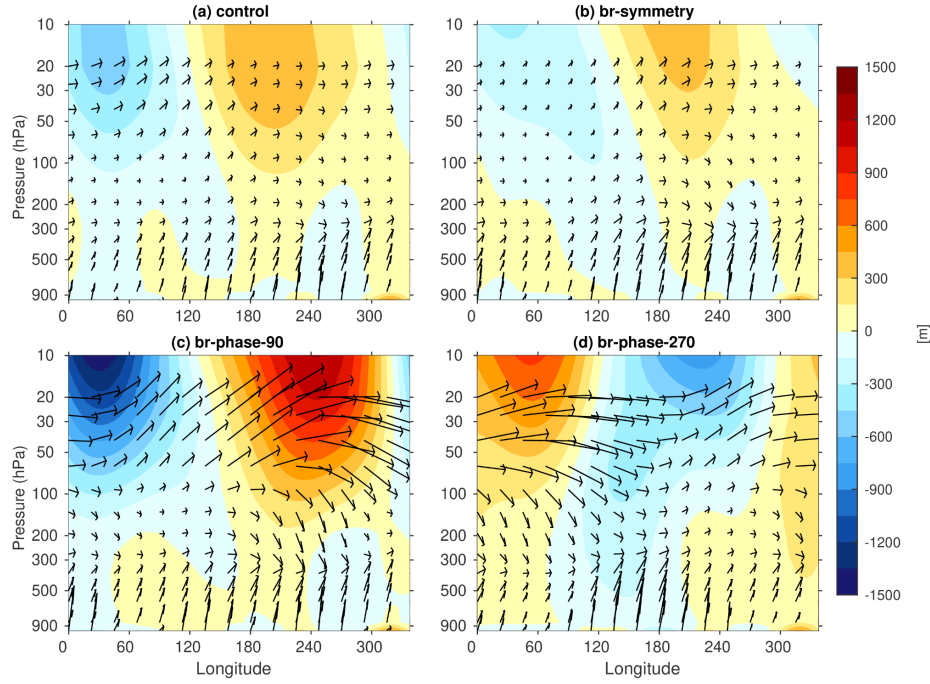


Figure R1. Longitude-pressure section of raw Fx and Fz (arrows; zonal and vertical components of Plumb flux) and raw geopotential height eddy (shadings; remove the zonal mean) area averaged between 55°N and 80°N in days 6-12.

Figure R1 shows the longitude-pressure section of raw Fx and Fz and raw geopotential height eddy area-averaged between 55°N and 80°N in days 6-12. The planetary waves propagate upward and are reflected back mainly over North America in the symmetry and phase-90 ensembles (Fig. R1 b,c) but dominantly over North Eurasia in the phase-270 ensemble (Fig. R1 d), which is consistent with the Fz in Figure 7 of the manuscript. In the wave reflection regions, the geopotential height eddy shows an eastward tilt, which also indicates a wave reflection event. We want to highlight the impact of zonal asymmetries of the polar vortex on the zonal asymmetries of wave reflection events. Therefore, we also show the map of Fz in this manuscript. Figure R1 above is now included in Figure 6 in the revised manuscript.

Zonal asymmetries of the wave reflection event are closely associated with the zonal asymmetries of the stratospheric polar vortex. We use the mass streamfunction to reveal the downward dynamical coupling from the stratosphere to the troposphere due to the additive effects from the polar vortex and wave reflection event, because they are dynamically coupled with each other.

We add some discussion on the limitations of MiMA starting from L132.

Note that wave reflection events could happen during both strong and weak polar vortex states, and we focus on the wave reflection event during SSWs in this study. Please see Table A1 in the corrigendum of Messori et al. (2022). It shows 45 selected stratospheric reflection events and clarifies 9 reflection events associated with SSWs from 1980 to 2021.

The authors claim to be the first to demonstrate that downward wave propagation can causally lead to surface impacts (L86, 427). I am concerned that this might be overstated; at least some of the aspects have been shown in the existing literature. Dunn-Sigouin and Shaw (2018) investigate the mechanism of extreme stratospheric wave-1 negative heat flux events with nudging experiments in a dry dynamical core model. They demonstrate the causal role of stratospheric waves in reproducing the tropospheric responses to negative eddy heat flux events. Through stratospheric nudging experiments in an AGCM, Ding et al. (2023) provide causal evidence that strong stratospheric wave activity leads to North American cold anomalies through vertical wave coupling. Their strong stratospheric wave events exhibit regional raw negative Plumb flux over North America (Fig. 2). They also analyzed weak wave events, during which the stratospheric planetary waves are in an opposite phase compared to those of strong wave events.

Neither of these studies presents exactly the same analysis/results as this study, so I think this analysis is still novel. But the authors should include these publications in their introduction/discussion. And perhaps the point can't be quite so strongly stated.

Thank you for sharing these valuable papers. We have carefully compared them with our results and cited these papers in the suitable places of the revised manuscript.

Individual comments

L5: The experiment names “90E” and “270E” do not feel very informative, as they represent the longitude where the zonally asymmetric component is 0. I'm unsure if it is a good idea, but the authors may consider using the longitude of the largest easterly forcing.

The location of the largest easterly forcing is not consistent with the center of row Z10 in Figure 2. Therefore, we remove this part “centered at 90E vs. 270E” from the abstract, just using the “phase-90” or “phase-270” to name the two experiments by specifying the longitudinal phase and directly linking them to equation 4.

L56: Huang et al. (2018) show that the shifting direction of displacement events affects the surface responses in reanalysis.

Cited near L32.

L58: An imposed heating induces an effect on meridional circulation opposite to what?

An imposed heating would induce a qualitatively opposite effect on the meridional circulation in the naturally occurring SSWs. In naturally occurring SSWs, the strengthened downwelling induces intensified adiabatic heating and results in stronger poleward meridional circulation in the stratosphere, while the imposed heating would instead produce upwelling and thereby equatorward meridional circulation in the stratosphere.

L66: Ding et al. (2022) show the distinction between zonally symmetric and asymmetric stratospheric variabilities.

Cited in L33.

L88: Could you point me to the reference that suggests wave reflection events often occur after SSW events? A single case of 2021 does not read like enough evidence.

Please see Table A1 in the corrigendum of Messori et al. (2022). It shows 45 selected stratospheric reflection events and clarifies 9 reflection events associated with SSWs from 1980 to 2021.

L139: easterly -> westerly

Corrected.

L141: The SSW frequency seems much lower than observed: 48 events in 9 runs of 50 years. How so?

We focus on only the JFM period, excluding November and December, which would give a lower ratio. White et al. (2022) reported a ratio of 0.29 SSWs per year in their control runs by using MiMA. This ratio is still smaller than the observed ~ 0.67 SSWs per year, while slightly different gravity wave settings lead to an SSW frequency above 0.4 per year.

This discussion has been added to L132 in the revised manuscript.

L171: Why set the scaling factor of the asymmetric component to 4? Do other settings change the conclusion?

We set the scaling factor to 4 to allow for a substantial zonally asymmetric component of the vortex response. Lowering it leads to a weaker regional signal, and strengthening it leads to a stronger regional signal.

We have also performed experiments with a wide range of zonally asymmetric scaling factors, though for all but the final configuration the model output wasn't retained, and plots weren't saved for most configurations either. But to give the reviewer a sense that results are insensitive, below we include a figure (Figure R2) showing the Z 500hpa anomalies in week 1 for

a scaling factor of 0.8 (the paper uses 4; Figure R3) for the zonally asymmetric component. The anomalies are very linear. A larger scaling factor amplifies regional anomalies, while the spatial pattern of responses is essentially identical in Figures R2 and R3. Note that the anomalies in Figures R2 and R3 are the differences between phase-90/270 and symmetry ensembles.

This discussion has been added to L181 in the revised manuscript.

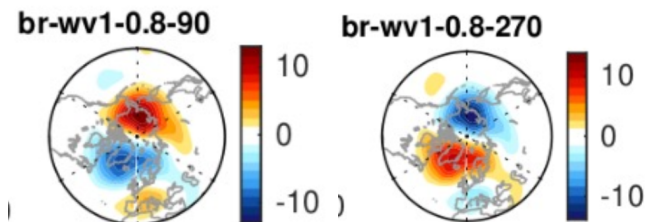


Figure R2. Geopotential height anomalies (m) at 500 hPa in week 1 with scaling factor 0.8. Note that the anomalies are the differences between phase-90/270 and symmetry ensembles.

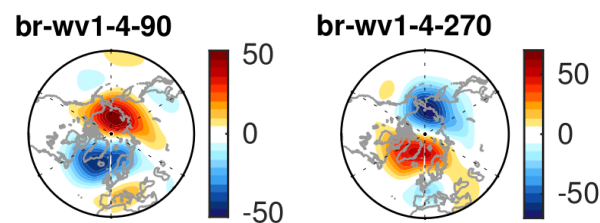


Figure R3. Geopotential height anomalies (m) at 500 hPa in week 1 with scaling factor 4. Note that the anomalies are the differences between phase-90/270 and symmetry ensembles.

Figure 2: Why show the geopotential height patterns during days 1 to 5 while showing zonal wind anomalies during days 6 to 12? Do the two periods differ a lot?

There are no big differences between days 1-5 and days 6-12, except for the more intensified anomalies in days 6-12. We have replaced the days 1-5 panels in Figure 2 with the geopotential height patterns during days 6 to 12 in the revised manuscript.

Figure 2: I suggest adding contours of 10 hPa anomalies, which would help readers understand through a different lens. Comparing the 10 hPa anomaly with the 500 hPa anomaly would provide insight into the vertical structure.

L245: This is confusing without showing the stratospheric anomalies.

Good suggestion, please see the new Figure 2 in the revised manuscript.

L256: The downward propagation seems hard to see, as the peak anomalies at 500 hPa occur at the same time as those at 10 hPa (day 12) for the branch experiments. In contrast, a clear lag

exists between stratospheric and tropospheric anomaly peaks for SSWs in the control experiment.

The external forcings are applied only above 100 hPa in the stratosphere. Consequently, the tropospheric anomalies arise from the downward coupling of the stratospheric perturbations. On shorter timescales (e.g., 6-hourly), the tropospheric response lags behind the stratospheric anomalies.

L265: There is no visible signal during days 1 to 12. How should we see the differences?

The cooling or warming signal is weak and not significant in days 1-12, but the spatial pattern is consistent with that occurring in days 13-33 for each branch ensemble, and the spatial pattern differs a lot between different ensembles, especially between phase-90 and phase-270 ensembles.

L283: Precipitation anomalies following symmetric forcing and control SSWs appear to be more persistent.

This makes sense. The composite of control SSWs shows a similar geopotential height pattern as the symmetry ensemble (as seen from Figure 2) because it includes many different types of SSWs with different zonal asymmetries. Here, we modify this sentence to be “precipitation anomalies weaken rapidly in phase-90 and phase-270 ensembles” in the revised manuscript.

L311, 410: Cluster 5 in the references appears to represent weak polar vortex events. Please elaborate on how they are connected to the downward wave propagation reported here.

Figure 2d of Kretschmer et al. (2018) shows the anomalous F_z at 100 hPa for cluster 5. Anomalous negative F_z emerges over both the east and west of North Eurasia, which indicates the wave reflection events there. Cohen et al. (2021) only presented the F_z at 100 hPa for cluster 4 in their Figure 3, although we discussed with Cohen the results of cluster 5 offline. This paper (Cohen et al. 2021) has been removed from here in the revised manuscript.

Figure 7: Why do the centers of poleward anomalies seem to shift 90 degrees westward from day 1-5 to day 6-12 in panels (c) and (d)?

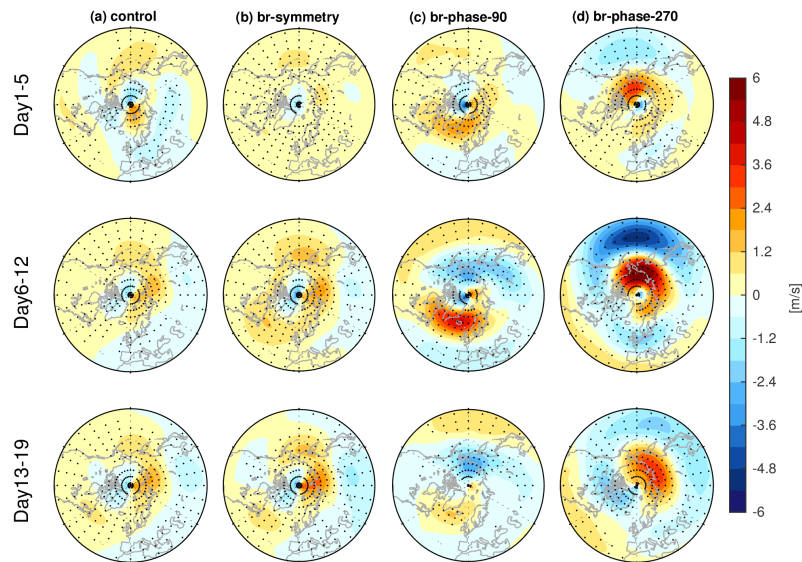


Figure R4. Anomalies of the divergent part of the meridional wind at 10 hPa for (a) control run, (b) symmetric branch ensemble, (c) phase-90 ensemble, and (d) phase-270 ensemble. Stippling indicates statistically significant anomalies at the 95% level based on Student's t test.

Figure R4 shows the anomalies of the divergent part of meridional winds at 10 hPa. In days 6-12, the positive anomalies of the divergent part of meridional wind are intensified in both the phase-90 and phase-270 ensembles (Figure R4 c,d; middle row), while the negative anomalies emerge in the sector opposite to the positive anomalies, resulting in anomalous downwelling there, thereby intensifying the anomalous high at 10 hPa (Figure 2 in the manuscript). Therefore, the shift of the center of poleward anomalies is due to the development of the polar vortex.

L363: What do you mean by “wave reflection during weak vortex events”? Perlwitz and Harnik (2003) show that a reflective surface is more likely to form when the lower stratospheric polar vortex is anomalously strong. Ding et al. (2022) also present a stronger-than-normal polar vortex prior to strong stratospheric wave events that are associated with regional wave reflection.

The wave reflection event could happen during both strong and weak polar vortex states, and we focus on the wave reflection event during SSWs in this study. Please see Table A1 in the corrigendum of Messori et al. (2022). It shows 45 selected stratospheric reflection events and clarifies 9 reflection events associated with SSWs from 1980 to 2021.

References

Judah Cohen et al. ,Linking Arctic variability and change with extreme winter weather in the United States.Science373,1116-1121(2021).DOI:10.1126/science.abi9167

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