

1 **PV power modelling using solar radiation from ground-based**
2 **measurements and CAMS: Assessing the diffuse component related**
3 **uncertainties leveraging the Global Solar Energy Estimator (GSEE)**

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24 **Abstract**

25 Accurate PV power production modelling requires precise knowledge of the distribution of solar
26 irradiance among its direct and diffuse components. Since this information is rarely available, this

27 requirement can be addressed through the use of diffuse fraction models. In this study, we try to
28 quantify the errors in PV modelling when measurements of the diffuse solar irradiance are not
29 available. For this purpose, we use total and diffuse solar irradiance data obtained from ground-
30 based measurements of BSRN to simulate the PV electric output using GSEE. We have chosen five
31 sites in Europe and North Africa, with different prevailing conditions, where BSRN measurements are
32 available. GSEE incorporates an implementation of the Boland-Ridley-Lauret (BRL) diffuse fraction
33 model, along with a Climate Data Interface that enables simulations across different time scales.
34 We evaluate the capability of BRL in providing accurate estimations of the diffuse fraction under
35 diverse atmospheric conditions, with particular attention on the presence of clouds and aerosols
36 and assess the extent to which its associated errors propagate to energy production modelling.
37 Furthermore, we compare GSEE outputs when using CAMS radiation time-series as input instead of
38 ground-based measurements, to quantify the impact of the CAMS radiation product uncertainties in
39 PV modelling.

40 **Keywords**

41 Solar energy modelling; CAMS radiation; PV power modelling; aerosol; dust; solar radiation

42 **1. Introduction**

43 Decarbonizing the power sector in a sustainable manner is pivotal in the effort to mitigate climate
44 change (Edenhofer et al., 2011; Owusu & Asumadu-Sarkodie, 2016; IPCC, 2023) and the large-scale
45 deployment of Solar Energy offers significant prospects toward this objective (Kakran et al., 2024).
46 The available solar energy is a variable source, fluctuating across different timescales with a unique
47 solar-resource profile over individual locations (McMahan et al., 2013). Therefore, accurate solar
48 energy forecasting and resource assessment is crucial for minimizing the risk in selecting project
49 location, designing the appropriate solar-energy conversion technology, and integrating new sources
50 of solar based power generation into the electricity grid (Stoffel, 2013), while short-term, intra-hour
51 forecasts are critical for power plant operations, grid-balancing, real-time unit dispatching,
52 automatic generation control, and trading (Pedro et al., 2017).

53 Extending solar irradiance forecasting to derive PV power forecasts is essential in solar energy
54 applications. PV power modelling can be achieved through the following additional steps to solar
55 irradiance forecasting: (i) decomposing Global Horizontal Irradiance (GHI) into Diffuse Horizontal
56 Irradiance (DHI) and Direct Normal Irradiance (DNI); (ii) calculating the plane-of-array irradiance

57 incident on the surface of PV planes, whether static or mounted on a solar tracking system, and (iii)
58 simulating the PV power production primarily based on the in-plane irradiance (Blanc et al., 2017).

59 The scarcity of concurrent measurements of both solar irradiance components, coupled with the
60 complexity of their theoretical computation, has driven the development of numerous empirical
61 models for estimating the diffuse fraction (ratio of the diffuse-to-global solar radiation). A seminal
62 contribution in this area was made by Liu and Jordan (1960), who established a correlation between
63 the diffuse fraction and the clearness or cloudiness index (ratio of the global-to-extraterrestrial
64 radiation). These models predominantly rely on the clearness index as the principal predictor. They
65 are generally classified into single-predictor models and multi-predictor models, with the latter
66 incorporating additional astronomical variables for enhanced precision (Paulescu & Blaga, 2019).
67 Typically, these models are expressed as polynomial equations, ranging from the 1st to the 4th degree,
68 that link the diffuse fraction to the clearness index $DF = f(\text{clearness index, } * \text{params})$ (Jacovides
69 et al., 2006). Boland et al. (2001) proposed the use of a logistic function instead of linear or simple
70 nonlinear functions of the clearness index. Ridley et al. (2010) developed a multiple-predictor
71 logistic model, known as the Boland-Ridley-Lauret (BRL), which combines simplicity and reliable
72 performance across both the Northern and Southern Hemispheres. The BRL model extends Boland's
73 approach by adopting the hourly clearness index as the principal predictor and introducing the
74 following additional parameters: apparent solar time, daily clearness index, solar altitude, and a
75 measure of the persistence of global radiation level. In the implementation of the BRL included in
76 the GSEE, the users set as input only the hourly clearness. Moreover, this implementation adopts the
77 updated parameters proposed by Lauret et al. (2013), which derived using data from nine worldwide
78 locations covering a variety of climates and environments across Europe, Africa, Australia and Asia.
79 While the existing models consider all-sky conditions, in solar energy modelling it is critical to focus
80 on cloud-free skies, where energy production is maximized. Under such conditions, aerosols
81 become the primary parameter influencing the distribution of solar irradiance among its
82 components. (e.g., Blaga et al., 2024). Specifically, the BRL model accounts for aerosols indirectly
83 through the clearness index, which is indicative of the overall atmospheric attenuation of solar
84 radiation.

85 In regions dominated by abundant sunshine, such as the Mediterranean and Middle East, which are
86 favorable for solar based power generation, the attenuation of solar irradiance is strongly influenced
87 by aerosols, and particularly desert dust aerosols. Several studies highlighted the impact of desert

88 dust aerosol in the downwelling solar irradiance and the energy production in these regions
89 (Fountoulakis et al., 2021; Papachristopoulou et al., 2022; Kosmopoulos et al., 2018; Kouklaki et al.,
90 2023). The significance of considering the effect of aerosols in short-term solar irradiance forecasting
91 and nowcasting is emphasized by Kazantzidis et al. (2017), Raptis et al. (2023) and
92 Papachristopoulou et al. (2024).

93 The Global Solar Energy Estimator (GSEE; Pfenninger & Staffell, 2016) is a widely used open access
94 model for simulating PV power output, designed for rapid calculations and ease of use. It comes with
95 an implementation of the BRL diffuse fraction model (Ridley et al., 2010; Lauret et al., 2013).

96 While PV power modelling is essential for linking solar resources to energy production, the existing
97 literature does not adequately address its reliability under diverse atmospheric conditions. To the
98 best of our knowledge, the existing literature does not include studies that explicitly address the
99 uncertainties in PV energy production modeling associated with the partitioning of solar radiation
100 into its direct and diffuse components at the model input. In this study, we supply GSEE with input
101 data from ground-based measurements as well as from the Copernicus Atmospheric Monitoring
102 Service (CAMS), aiming to investigate differences in PV power output simulations, which arise from
103 providing only GHI as input radiation data. At the outset, we focus on evaluating the reliability of BRL
104 under diverse atmospheric conditions, with particular attention to the dependence of its accuracy
105 on the presence of clouds and aerosols. To further explore this, we conduct a sensitivity analysis
106 using radiative transfer model (RTM) simulations under cloud-free skies. Following these analyses,
107 we assess the extent to which the associated uncertainties in the estimation of the diffuse fraction
108 spread to the power generation over hourly intervals. This step involves simulating PV plants with
109 varying configurations. GSEE is also effective for analyzing trends and variability in solar based power
110 generation through its climate interface submodule (e.g., Hou et al., 2021), where the BRL model is
111 integrated within the internal processing chain. The accuracy of the climate interface in estimating
112 the total daily PV power output is also evaluated in this study.

113

114 **2. Data and Methodology**

115 2.1 Global Solar Energy Estimator (GSEE)

116 The modelling of the PV power output is conducted using the version 0.3.1 of GSEE (Pfenninger &
117 Staffell, 2016). The model features functions for simulating a complete PV system, incorporating

118 characteristics and specifications such as location, installed capacity, technology, tracking (fixed, 1-
119 axis, 2-axis), tilt angle, and orientation.

120 The user provides as input time-series data of solar radiation, and optionally, ambient air
121 temperature and surface albedo. Specifically, the model requires GHI and, when available, the
122 Diffuse Fraction. If the diffuse component is not provided, the provided implementation of the BRL
123 diffuse fraction model (Ridley et al., 2010; Lauret et al., 2013) is employed to estimate it, relying only
124 on time-series of the hourly clearness index and the geographical coordinates. While in the single-
125 site application of the GSEE model with hourly time resolution the user has the option to adjust the
126 input and select alternative diffuse fraction models implemented by external libraries, e.g., pvlib
127 (Anderson et al., 2013), the climate data interface automatically invokes the BRL model as part of the
128 internal processing workflow. GSEE utilizes the provided information for the distribution of the
129 irradiance components and applies trigonometric calculations to determine the total solar
130 irradiance incident on the panel's inclined plane. More precisely, for the plane-of -array irradiance
131 calculation a GSEE includes the submodule "trigon" (transposition model), which is based on
132 trigonometric formulations, that account of the surface albedo, thereby including the ground-
133 reflected component of solar radiation. However, the transposition model is integrated within the
134 GSEE internal algorithms, so it cannot be modified by the user.

135 After solar irradiance the most significant parameter regarding energy production is air temperature
136 (e.g., Dubey et al., 2013). If temperature is not provided by the user, the model assumes a default
137 value of 20 °C. In this study, temperature was used as input only in the simulations with BSRN data,
138 as it is provided alongside radiation measurements. A surface albedo value of 0.3 considered by
139 default from the model, introduces some uncertainty in our simulations, which however is estimated
140 to be small. Under cloudless conditions, a 10% difference in surface albedo changes the GHI by ~1%
141 for $SZA < 75^\circ$. Differences are larger under cloudy conditions (~ 10% difference in GHI for a 10%
142 difference in surface albedo). Nevertheless, surface albedo at the selected sites is generally low and
143 relatively invariant throughout the year (even at the most northern site of Lindenberg there is only a
144 limited number of days with increased surface albedo due to snow cover).

145 The available options for the panel type are crystalline silicon (c-Si) and Cadmium Telluride (CdTe),
146 where the power output is modeled based on the relative PV performance model described by Huld
147 et al. (2010). For fixed panels, a built-in latitude dependent function for the optimal tilt is also
148 included.

149 Moreover, GSEE includes a Climate Data Interface submodule that enables the processing of gridded
150 climate datasets, with varying temporal resolutions, ranging from hourly to annual. Within the
151 context of this submodule, the use of BRL serves as part of the resampling and upsampling
152 processes applied to input climate datasets with daily resolution. For processing data with lower-
153 than-daily resolutions, it incorporates the use of Probability Density Functions (PDFs), which
154 describe the probability with which a day with a certain amount of radiation occurs within a month
155 (Renewables Ninja, n.d.). This methodology accounts for the non-linear distribution of mean monthly
156 radiation across individual days, ensuring a more representative temporal disaggregation. The
157 processes applied to the mean daily irradiance are described in detail in Section 3.4.

158 For the purposes of this study, we simulated solar plants with capacity of 1 kWp, and for both
159 available technologies. The simulations with c-Si technology, considered as default by the model,
160 are presented in detail the following sections. The results of the simulations with CdTe technology
161 are provided in the supplement, and are not thoroughly discussed, since they are very similar to the
162 results for the c-Si technology. Regarding the mounting approach, the solar plants were either static
163 and oriented to the south or equipped with a 2-axis solar tracking system. In the case of fixed panels,
164 we selected the optimal tilt angle relying on the latitude dependent built-in function.

165 The input parameters defining the characteristics of the simulated PV plants are summarized in Table
166 1.

167 **Table 1.** Input parameters defining the characteristics of the simulated PV plants

Capacity	Mounting Approach			Technology	
1 kWp	Fixed		2-axis tracking	c-Si	CdTe
	Orientation: south	Tilt Angle: $f(\text{latitude})$ built-in function for optimal tilt			

168

169 2.2 Ground-based measurements

170 We supplied GSEE with ground-based irradiance as well as ambient temperature measurements
171 collected from five stations of the Baseline Surface Radiation Network (BSRN; Driemel et al., 2018).
172 Moreover, information about aerosols was retrieved from co-located stations of the Aerosol Robotic
173 Network (AERONET; Holben et al., 1998; Dubovik et al., 2000).

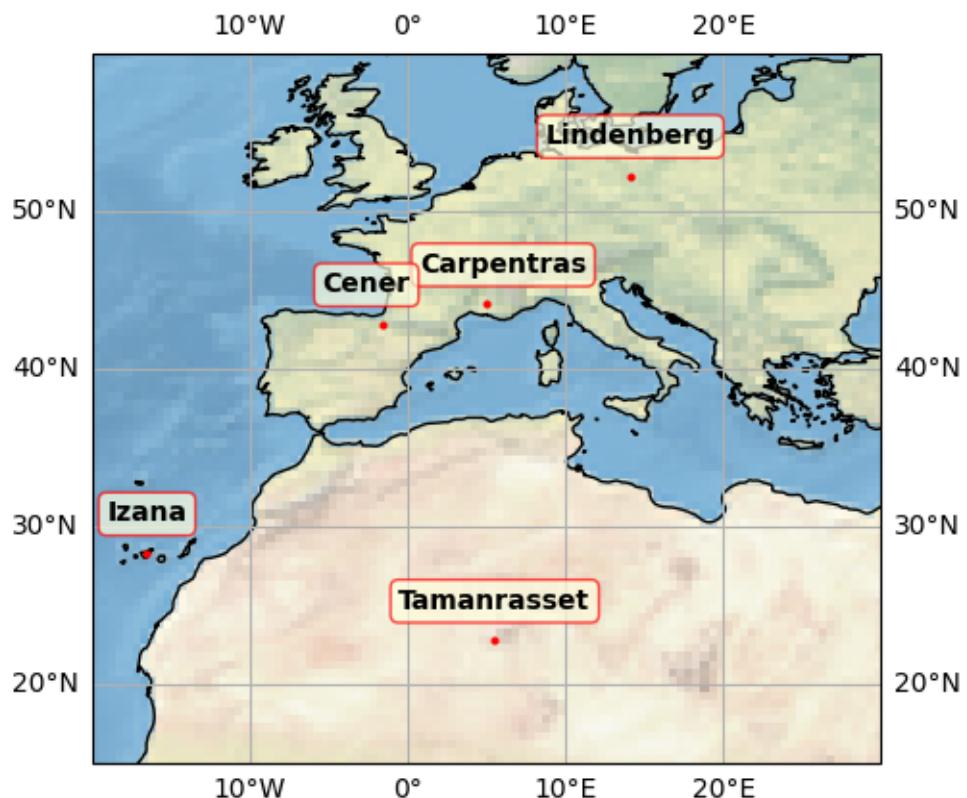
174 Information for the stations utilized for this study is summarized in Table 2, and their geographical
175 location is depicted in Figure 1.

176

177 **Table 2.** Detailed information about the location of the ground-based stations used in this study.

STATION	Latitude [° N]	Longitude [° E]	Elevation [m]
Carpentras (CAR)	44.08	5.06	100
Cener (CNR)	42.82	-1.60	471
Izaña (IZA)	28.31	-16.50	2373
Lindenberg (LIN)	52.21	14.12	125
Tamanrasset (TAM)	22.79	5.53	1385

178



179

180 **Figure 1.** Locations of the BSRN and co-located AERONET stations that are used in the current
181 study

182

183 BSRN station-to-archive files were accessed and manipulated using the SolarData v1.1 R package
184 (Yang, 2019), and the BSRN-recommended quality check (QC) tests (Long & Dutton, 2010) applied
185 to the collected data. Some data gaps arose due to measurements removed during the QC
186 procedure. Although these data gaps are, in most cases, shorter than 2-3 hours, they may affect the
187 BRL performance throughout the corresponding days. Consequently, days affected by such data
188 gaps excluded from the analysis. We retrieved data for 2017, with 1-minute temporal resolution. We
189 used GHI, DHI, and Temperature as inputs to the GSEE model. Initially, the data were resampled to
190 hourly and mean hourly values of GHI and DHI are calculated. Then, the simulations were conducted
191 using either GHI and DHI, or only GHI along with the deployment of BRL. The input to BRL consists of
192 hourly clearness index, derived by dividing GHI measurements with the solar radiation incident on a
193 horizontal plane at the Top of the Atmosphere (TOA) above the examined location. Subsequently, the
194 1-min timeseries resampled also to a daily resolution and transformed into three-dimensional
195 arrays, $GHI = f(\text{time}, \text{lat}, \text{lon})$, where the spatial dimensions of each dataset corresponded to a
196 unique point defined by the coordinates of the associated station. Simulations with the daily time-
197 resolved dataset were performed using the Climate Data Interface.

198 Representing cloudiness is a challenging task that requires several observations. For this purpose,
199 aiming to obtain an indicative measure of the intra-hour cloudiness conditions we adopted the
200 following formulation. Specifically, measurements of Direct Normal Irradiance (DNI) were utilized to
201 obtain information for cloudiness relying on the conditions stated by WMO (2021), according to
202 which sunshine duration is the total period where DNI exceeds 120 W/m^2 . Alternative approaches
203 such as the Cloud Modification Factor, require estimates of the clear sky irradiance, which
204 introduces additional uncertainty. For the purpose of this analysis, we introduced a solar visibility
205 (SV) parameter. Specifically, we assigned the value 0 when sun was obscured and the value 1 when
206 visible. Aiming to describe the mean intra-hour cloudiness conditions, we considered the sky as
207 cloud-free, cloudy, and partly cloudy based on the mean SV for the entire corresponding hour as
208 follows:

$$\langle SV \rangle_{hour}: \begin{cases} 1 & \text{cloud-free} \\ \in (0,1) & \text{partly cloudy} \\ 0 & \text{cloudy} \end{cases}$$

210 For aerosol information, we accessed the AERONET Version 3 (V3) (Giles et al., 2019) and retrieved
 211 level 2.0 data (from direct sun measurements) for Aerosol Optical Depth at 500nm (AOD_{500}), which
 212 serves as a representative measure of the aerosol load; Ångström Exponent between 440 and 870
 213 nm wavelengths ($AE_{440-870}$), where values near 0 correspond to coarse dust particles and values
 214 around 2 to fine (e.g., smoke) particles (Dubovik et al., 2002); and Fine Mode Fraction at 500nm
 215 (FMF_{500}) obtained from the Spectral Deconvolution Algorithm (SDA) retrievals, to distinguish aerosol
 216 into fine and coarse mode. The data were resampled at hourly intervals and a mean hourly value
 217 calculated. After, the hourly mean values divided into clusters based on AOD_{500} , reflecting different
 218 levels of aerosol load and allowing us to quantify their impact on solar energy production. To
 219 investigate the impact related exclusively to aerosols, we included only hours with cloud-free sky
 220 conditions. The clusters are defined in detail as follows:

- 221 • $AOD_{500} \leq 0.05$: Low aerosol load
- 222 • $0.05 < AOD_{500} \leq 0.15$: Moderate aerosol load
- 223 • $0.15 < AOD_{500} \leq 0.3$: High aerosol load
- 224 • $AOD_{500} > 0.3$: Very high aerosol load

225 To evaluate the performance of the Climate Interface over daily intervals, we defined the sunny
 226 (cloudless) days using the condition: $\langle SV \rangle_{day} \geq 0.9$. Next, to characterize the average aerosol
 227 conditions on sunny days, we applied the following classification:

- 228 • $\langle AOD_{500} \rangle_{day} \leq 0.05$: very-low aerosol
- 229 • $\langle AOD_{500} \rangle_{day} > 0.05$: aerosol-laden

230 Detailed comparisons of the energy production over hourly and daily integrals under the various
 231 predefined sky conditions are provided in the supplement through evaluation metrics.

232 The selected locations have quite different atmospheric conditions regarding cloudiness and
 233 aerosols. Additionally, they vary in altitude. A brief overview of the prevailing conditions derived from
 234 the ground-based data is provided on the supplement. Regarding cloudiness, it is notable that in
 235 Lindenberg the sky is generally overcast, whereas in southern locations sunshine dominates. In
 236 terms of aerosols, very high aerosol loads occur more frequently in Tamanrasset. As for aerosol type,

237 there is considerable variation among the examined locations: Carpentras, Cener, and Lindenberg
238 are primarily influenced by fine mode aerosols, while Tamanrasset and Izaña are mostly affected by
239 coarse mode aerosols.

240 For investigating the impact of desert dust aerosol in solar based power generation, Tamanrasset
241 serves as a representative and exceptional case because it is in a region with important sources of
242 Saharan dust aerosols (Faid et al., 2012). Meanwhile, Izaña, located in subtropical North Atlantic, is
243 a high mountain station within the free troposphere, affected mineral dust when the Saharan Air
244 Layer top exceeds the station height, especially through August to October (Toledano et al., 2018;
245 Cuevas et al., 2018). Due to its high altitude, Izaña avoids contamination from local or regional
246 sources (Barreto et al. 2022). The Canary Islands, where Izaña is located, are influenced by extreme
247 dust events that cause a significant decrease in PV power generation (Canadillas-Ramallo et al.,
248 2021). In South Europe, which is also affected by the transport of Saharan dust across the
249 Mediterranean, aerosol types exhibit a mixture as a result of simultaneous local pollution and low
250 concentration of mineral dust (Logothetis et al., 2020).

251 2.3 Copernicus Atmospheric Monitoring Service (CAMS)

252 We retrieved data from the CAMS radiation service (Schroedter-Homscheidt et al., 2022; Qu et al.,
253 2017), from the solar radiation time-series product (CAMS, 2020). The CAMS solar radiation service
254 provides historical estimates for global solar radiation, along with its components, from 2004 to
255 present. These values are provided with a frequency as fine as 1-minute. In this study, we used the
256 hourly time-series of GHI and DHI for all-sky conditions, setting the input coordinates to match the
257 locations of the BSRN stations. The solar radiation time-series product (CAMS, 2020) performs
258 interpolations integrated in its internal algorithm and provides time-series for the coordinates and
259 the altitude of a single-site location. We compared the solar energy production derived from the use
260 of CAMS data with that derived from the use of ground-based measurements from BSRN.

261 2.4 Radiative Transfer Model (RTM)

262 We performed Radiative Transfer (RT) simulations aiming to further assess the uncertainties in
263 estimating the diffuse fraction arising from the effect of aerosols. The simulations were conducted
264 using libRadtran (Emde et al., 2016; Mayer & Kylling, 2005), a widely used software package, allowing
265 the computation of radiances, irradiances, and actinic fluxes. A sensitivity analysis was performed
266 by comparing the diffuse irradiance calculated from libRadtran with the estimations of BRL. This

267 analysis examines the dependence of the aerosol-related discrepancy as function of Solar Zenith
268 Angle (SZA) and latitude, considering the effect of parameters such as surface albedo and altitude.
269 To conduct aerosol parameterizations, we considered the default aerosol extinction profile (Shettle,
270 1989) and set asymmetry factor (gg) to 0.7, while varying the Single Scattering Albedo (SSA) and the
271 Ångström Exponent (AE), and defining AOD_{500} by adjusting the value of the parameter-b in
272 Ångström's law (Ångström, 1929) as follows:

273
$$\tau_\lambda = b \cdot \lambda^{-a} \rightarrow AOD_{500} = b \cdot (0.5 \mu m)^{-AE}$$

274 The standard aerosol profiles (Anderson et al., 1986) were used for all sites. According to
275 Fountoulakis et al. (2022), using a more accurate vertical distribution of aerosols in the troposphere
276 would have a negligible effect in the GHI and DHI at the Earth's surface.

277 Table 3 illustrates the libRadtran settings used in this study.

278 **Table 3.** LibRadtran inputs

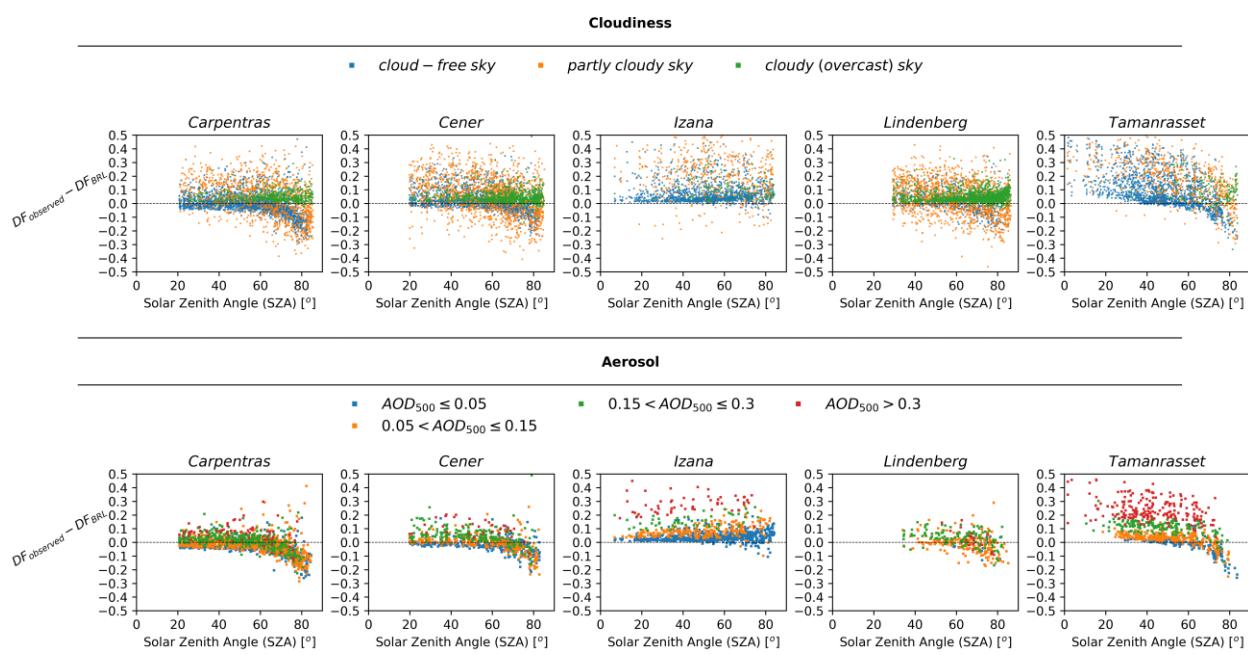
Parameter	Input
Atmospheric profile	Mid-latitude summer (April-September)/mid-latitude winter (October - March) (Anderson et al., 1986)
Extraterrestrial spectrum	(Kato et al. 1999)
Datetime	date and time input accompanied by project location coordinates
Altitude	0.1/2 km
Surface albedo	0.2 / 0.8
Number of streams	6
RT solver	sdisort (Buras et al., 2011)
AE	0 – 2 with step 1
SSA	0.7, 0.9, 1.0
gg	0.7
TOC (Total Ozone Column)	300 DU
Integrated Water Vapor	15 mm

279

280 **3. Results**281 3.1 Performance verification of the BRL diffuse fraction model

282 The performance of BRL was evaluated by comparing the actual diffuse fraction, obtained directly
283 from resampled to hourly BSRN ground-based measurements, with that derived using BRL. As a first
284 step, to isolate the influence of SZA from that associated with the atmospheric conditions, the
285 difference in diffuse fraction (DF) between the observed and the one estimated using BRL as a
286 function of SZA is presented in Figure 2. The atmospheric conditions are represented separately for
287 both all-sky and cloud-free sky conditions and are grouped into clusters, as outlined in Section 2.2.
288 The patterns reflecting the differences under the distinct sky conditions indicate an additional
289 dependency on SZA, which becomes apparent approximately at SZA between 60° and 70°. In most
290 cases, there is an almost constant displacement with respect to $y=0$ below 60°, as well as a change
291 in behavior when SZA exceeds this value. Izaña presents a special case, as the station is located at
292 a very high altitude. At such high altitudes the contribution of the diffuse component to the total
293 irradiance is significantly smaller relative to lower altitude sites, which seems to be captured more
294 accurately by BRL at high SZAs. We must also note that (i) at Izaña, the actual diffuse irradiance may
295 experience an additional enhancement due to the contribution of adjacent lower-lying clouds – an
296 effect that is not accounted for in the diffuse fraction model, and (ii) during dust events the site is
297 usually inside – and not under – the dust layer, which results in more complex interactions between
298 dust and solar radiation relative to lower altitude sites. Defining an exact limit (for the lower altitude
299 sites), where the behavior is changing, is challenging; therefore, 60° was selected for practical
300 energy-related applications, focusing on periods with meaningful energy contribution, and is
301 supported by the sensitivity analysis (Section 3.2) under clear-sky conditions. Concerning the same
302 grouped atmospheric conditions, Figure 3 illustrates the comparison between the observed and the
303 estimated diffuse fraction for $SZA \leq 60^\circ$. This approach allows us to examine BRL performance after
304 eliminating the influence of SZA, thereby providing a more comprehensive view of its reliability.

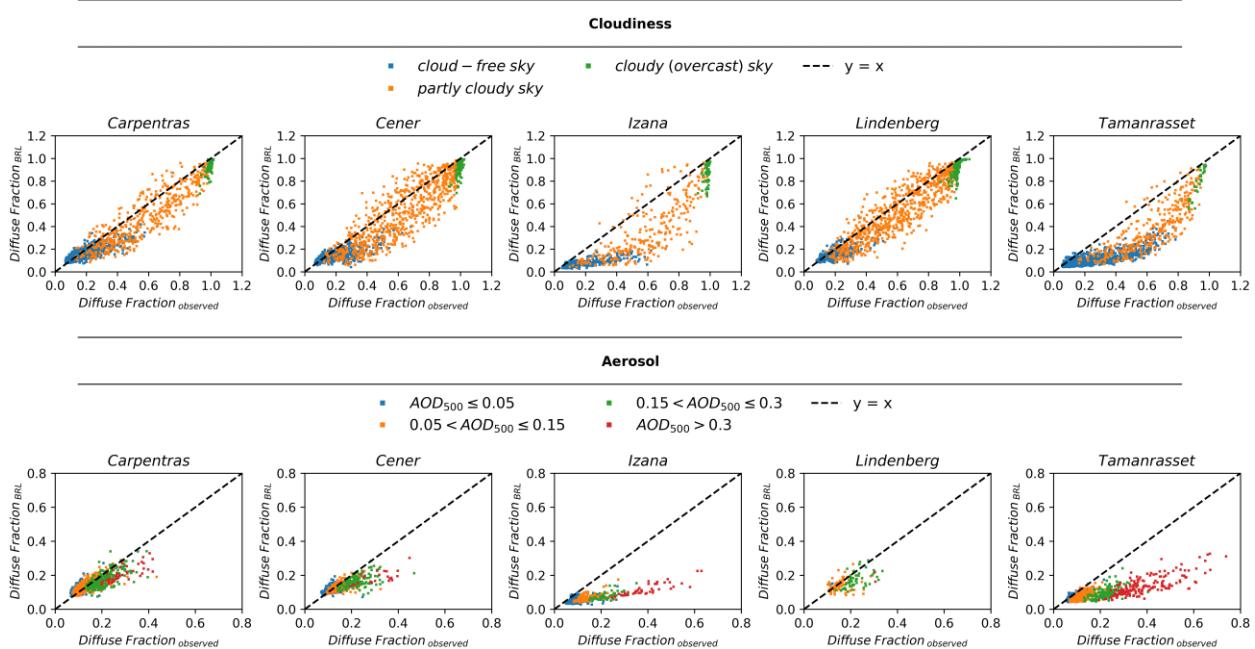
$$\text{Diffuse Fraction (DF) observed} - \text{Diffuse Fraction (DF)}_{\text{BRL}} = f(\text{sza})$$



305

306 **Figure 2.** Difference between the diffuse fraction estimated by the ground-based measurements
 307 and by using the BRL model as a function of SZA under diverse atmospheric conditions: (top)
 308 classification with respect to cloudiness and (bottom) classification with respect to aerosol optical
 309 depth

$$\text{Diffuse Fraction (DF)}_{\text{BRL}} = f(\text{Diffuse Fraction (DF)}_{\text{observed}})$$



310

311 **Figure 3.** Comparison of the diffuse fraction estimated using BRL with that estimated by the
312 ground-based measurements under diverse atmospheric conditions for $SZA < 60^\circ$: (top)
313 classification with respect to cloudiness and (bottom) classification with respect to aerosol optical
314 depth

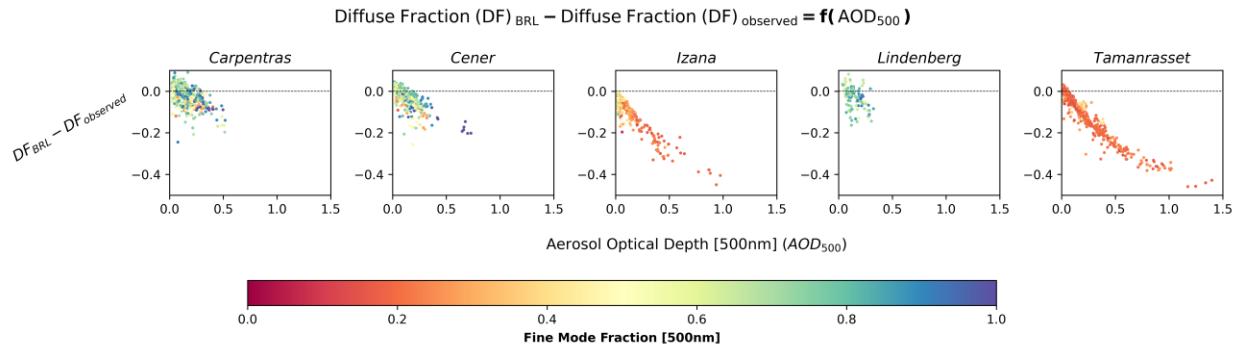
315

316 From Figure 3, a distinct dependency of BRL's reliability on the atmospheric conditions can be
317 observed. Under all-sky conditions, the presence of clouds has a notable impact on the model's
318 performance. Partly cloudy conditions result in greater dispersion of the values from the identity line
319 respectively, likely due to the complexity of such sky scenes. Under overcast conditions, where the
320 sky can be considered homogeneous and isotropic, the model in most cases performs slightly better.
321 However, the limitations of the DNI-based classification methodology, related to the complexity of
322 the cloud scenes, the spatiotemporal variability during the hourly periods, and the 3D variability of
323 cloud properties, would require additional observational tools for a more detailed investigation. More
324 specifically, the vast majority of overcast cases where the BRL diffuse fraction is below 0.8 while the
325 observed is close to 1 correspond to periods involving rapid transitions between partly cloudy and
326 overcast skies, occurring either during the hour itself or immediately before or after it. Furthermore,
327 a limited number of cases identified during intense dust events at Tamanrasset and Izaña, where the
328 reduction of DNI was so pronounced that the applied DNI-based criterion classified these conditions
329 as overcast. However, these cases are not further investigated, as the energy production levels during
330 such periods are very low.

331 Under cloud-free skies, BRL tends to underestimate, and this bias becomes more pronounced as
332 aerosol load increases. Aiming to highlight this dependency, Figure 4 shows the difference between
333 the estimated and the observed diffuse fraction as function of AOD_{500} , emphasizing also the extent
334 to which it is related to the aerosol type by providing FMF_{500} . A decrease for increasing AOD_{500} is
335 evident across all cases. In Tamanrasset and Izaña, associated with the influence of Saharan dust,
336 the coarse mode dominates, and a more distinct and well-defined curve is depicted compared to
337 other sites.

338 It is important to clarify that for assessing the impact of aerosols we have assumed entirely cloud-
339 free conditions. However, the criterion applied based on DNI does not fully guarantee the absence of
340 small, scattered clouds within the sky dome. Such clouds could induce slight enhancements in DHI.

341 A more rigorous assessment of the impact associated exclusively with aerosols could be achieved
 342 by integrating images from ground-based co-located all-sky cameras. On the other hand, the
 343 presence of aerosols even under cloudy scenes, introduces an additional uncertainty which is
 344 difficult to investigate accurately.



345
 346 **Figure 4.** Difference between the estimated using BRL and the diffuse fraction estimated by the
 347 ground-based measurements as function of AOD_{500} and FMF_{500}

348
 349 3.2 Sensitivity analysis of the BRL performance under cloud-free sky conditions from RT
 350 simulations
 351 The uncertainties in estimating diffuse fraction under cloud-free sky conditions, as discussed in
 352 section 3.1, are further investigated. We performed RT simulations using libRadtran to calculate GHI
 353 and DHI under various aerosol scenarios. The resulting GHI values were then used as input to BRL to
 354 estimate the diffuse fraction, which was subsequently compared to the diffuse fraction derived
 355 directly from the ratio of DHI to GHI computed by libRadtran.

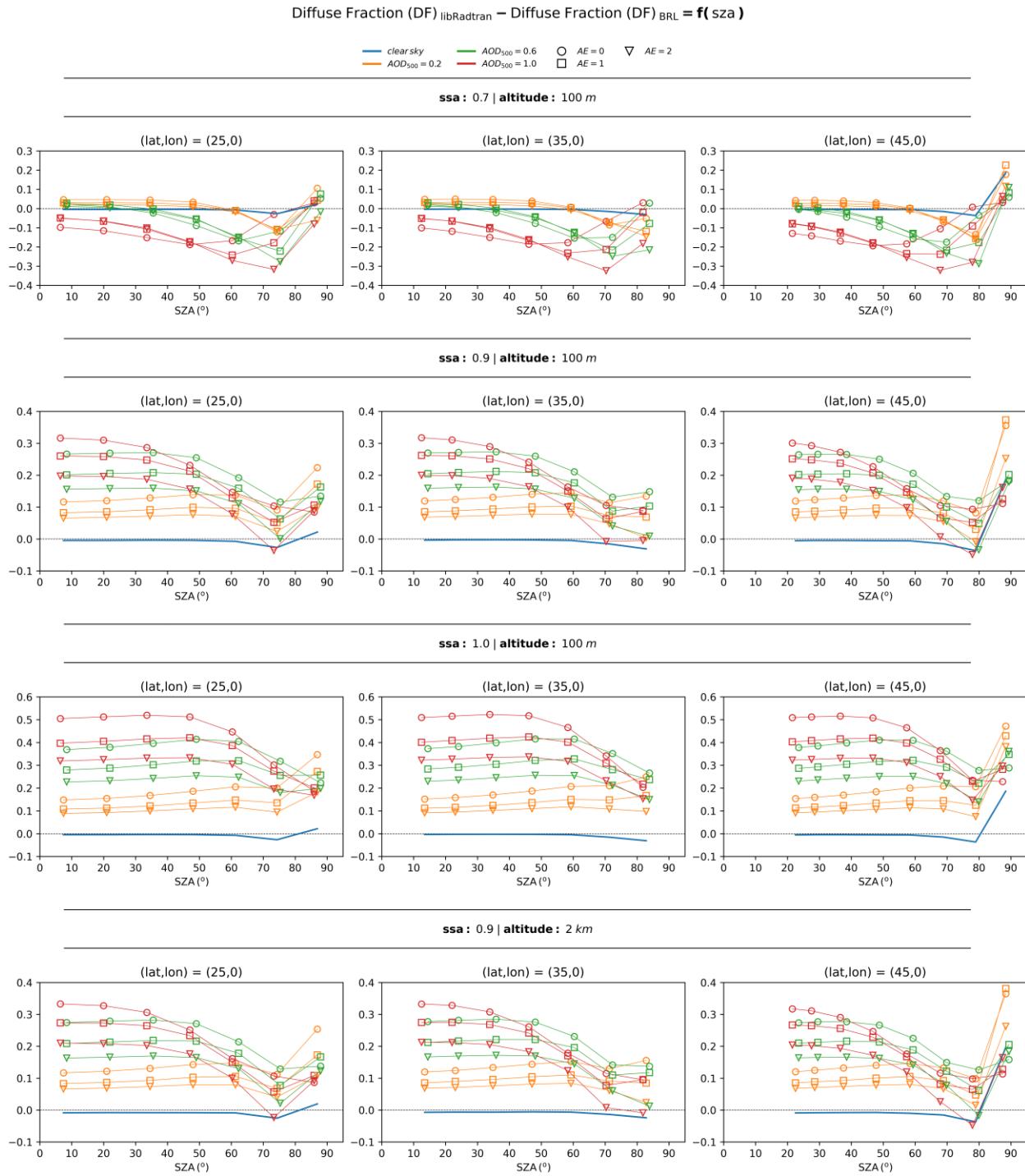
356 To ensure a comprehensive analysis, we considered three representative latitudes (25° , 35° and 45°).
 357 Since BRL requires an hourly time-series of GHI as input, the analysis was conducted for the summer
 358 solstice. On this day, a sufficient number of hourly values are available, corresponding to a wide
 359 range of SZA values, allowing for a robust assessment of the methodology. The sensitivity analysis
 360 was performed for surface albedo values of 0.2 and 0.8 as well as for altitudes of 0.1 and 2 km. For
 361 aerosol parameterization, we examined completely clear-sky conditions as a reference, alongside
 362 scenarios with AOD_{500} values of 0.2, 0.6, and 1, while varying the SSA and AE. Specifically, the
 363 scenarios included SSA values of 0.7, 0.9 and 1, combined with AE values of 0, 1 and 2. The results

364 of this sensitivity analysis for an albedo of 0.2 are provided in Figure 5, while the results for an albedo
365 of 0.8 are included in the supplement (Figure S1).

366 The results confirm that BRL performs well under clear sky conditions and for SZA below 60°, while
367 the incorporation of aerosols in the sky scene introduces larger uncertainties. In all scenarios, we
368 observe that lower values of AE correspond to higher uncertainties. Moreover, when SSA is 0.9 or 1
369 BRL gradually tends to underestimate the diffuse fraction as aerosol load increases. Instead, when
370 SSA is 0.7, BRL exhibits a different behavior, shifting toward an overestimation of the diffuse fraction
371 at high aerosol loads.

372 The findings of this sensitivity analysis are consistent with the evaluated BRL performance from
373 ground-based measurements presented in section 3.1, especially at SZA smaller than 60° - 70°, and
374 underscore the role of aerosol in the accuracy of diffuse fraction estimations. Differences between
375 the results shown in Figures 2 and 5 at SZA between 60° - 80° can be due to a number of site-related
376 reasons. For example, enhancement of the diffuse component due to scattering by underlying
377 atmospheric layers and clouds in the case of Izaña may compensate the observed overestimation of
378 the diffuse fraction by BRL. Concerning the impact related to AE and SSA, we confirm that the higher
379 underestimations observed for Tamanrasset and Izaña are associated with the optical properties of
380 desert dust aerosol particles. While AE and SSA alone are not sufficient to fully characterize the
381 aerosol type, they serve as strong indicators, aligning with the classification framework of Dubovik et
382 al. (2002). The same comparison for albedo 0.8 (Figure S1 in the supplement) reveals a significant
383 broadening of the discrepancies. Moreover, we observe the presence of a systematic error, even
384 under clear sky conditions.

385 The resulting differences were practically identical across the three selected latitudes, indicating
386 that the BRL model is largely independent of latitude and can therefore be considered as a reliable
387 solution over a wide range of latitudes. Furthermore, the effect of altitude was found to be small.
388 Finally, the outcomes of this analysis highlight potential inconsistencies arising from aerosols with
389 different optical properties. Although the updated parameters of the BRL's model (as implemented
390 in the GSEE model) reported by Lauret et al. (2013) were derived using data from nine worldwide
391 locations, encompassing a broad range of sky conditions that capture a fully representative set of
392 optical properties remain challenging.



393

394 **Figure 5.** Difference between the diffuse fraction derived directly from the computations of DHI and
 395 GHI using libRadtran and the one estimated by applying BRL to the libRadtran-computed GHI

396

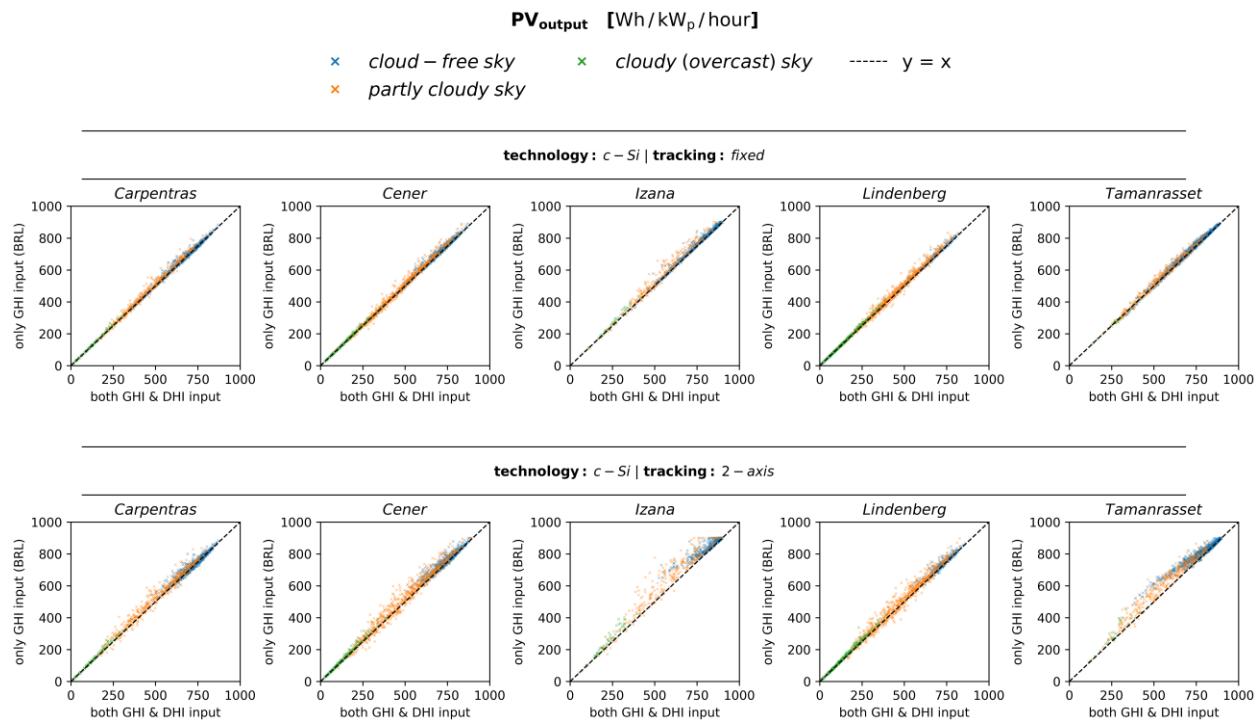
397

398 3.3 Analysis of the differences in energy production using hourly integrals within the modelling of PV
399 plants

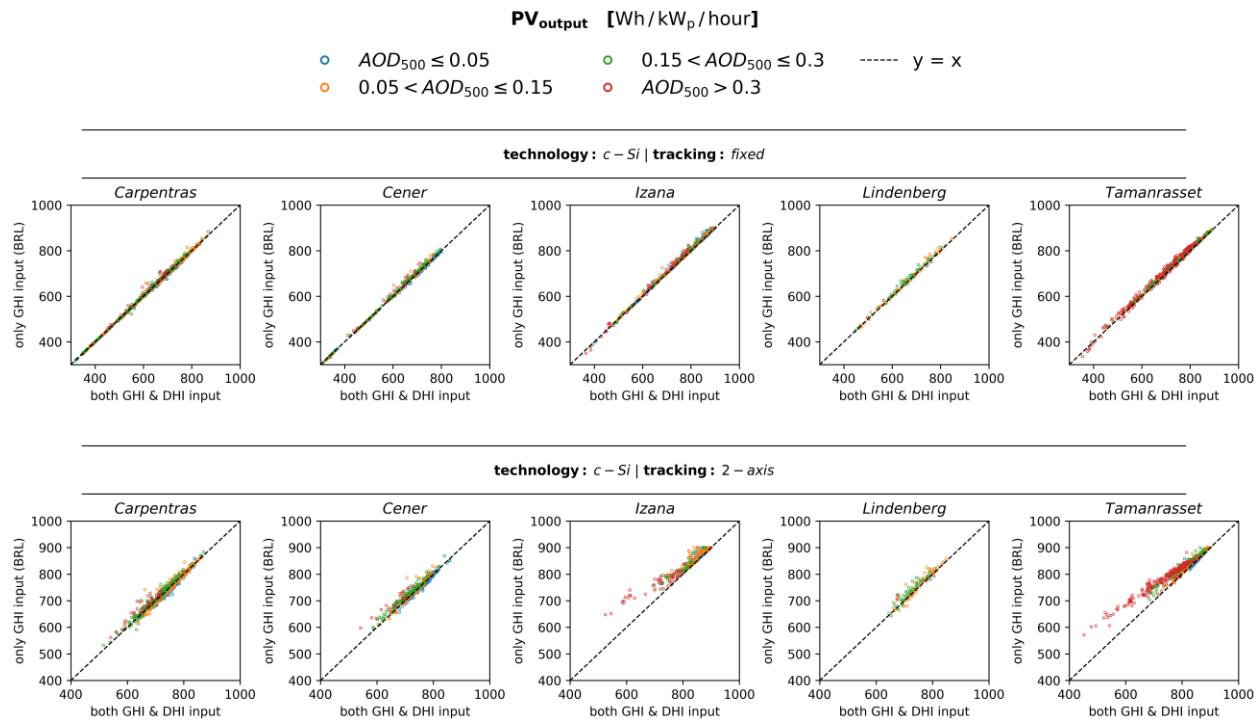
400 Uncertainties in estimating the diffuse fraction influence the calculation of the total irradiance
401 received by an inclined panel's surface, thereby affecting the accuracy of the PV power simulations.
402 In this section, we employ the main submodule of GSEE, used for modelling the electric output from
403 a PV panel, aiming to assess the extent to which these uncertainties propagate to the estimation of
404 the hourly power production. We analyze discrepancies arising from using only GHI from BSRN as
405 input radiation data to the model, instead of both DHI and GHI. More specifically, we compare the
406 total energy produced per hour per unit, expressed in watt-hours (Wh), per unit of nominal power
407 (kWp). The energy production is evaluated for both fixed panels and 2-axis tracking systems.

408 The results of this comparison for c-Si based technology PV panels for different atmospheric
409 conditions are presented in Figure 6, illustrating the impact of cloudiness, and in Figure 7,
410 demonstrating the effect of aerosols. The corresponding results for CdTe technology are provided in
411 the supplement (Figures S2 and S3 respectively). In the modelling of 2-axis solar tracking systems,
412 where the panel is continuously adjusted to maintain a perpendicular orientation to incoming solar
413 radiation, the system becomes more sensitive to uncertainties in the estimation of the diffuse
414 fraction, leading to more significant differences in energy production. Specifically, the contribution
415 of the direct irradiance is maximized in such systems, as the panel exploits the entirety of the
416 available direct irradiance. On the other hand, in the simulation of static panels, the contributions of
417 direct and diffuse components are more evenly distributed, making the impact of diffuse fraction
418 uncertainties less pronounced in energy production.

419 Regarding the uncertainties related to the atmospheric conditions, from Figure 6 we confirm that the
420 highest dispersion occurs in partly cloudy conditions, while from Figure 7, where we examine cloud-
421 free conditions, we note that further improvement achieved as aerosol load decreases. Under totally
422 overcast skies the energy production is extremely low, rendering errors practically negligible.
423 Moreover, accuracy is influenced by aerosols, where a gradual decline in accuracy is detected as
424 aerosol load increases. However, assessing the extent of aerosol loading impact is complex,
425 depending on the interaction of solar radiation with particles of varying optical properties, as
426 extensively analyzed in the previous sections. This effect becomes particularly evident in cases of
427 high aerosol loading, where a noticeable offset is observed, while under certain conditions, the
428 associated uncertainty is comparable to that found in partly cloudy conditions.



430 **Figure 6.** Comparison of the estimated hourly PV power generation between simulations performed
 431 using GSEE with input data consisting of either only GHI or both GHI and DHI under varying
 432 cloudiness conditions: (top) fixed panels (bottom) 2-axis tracking systems



434 **Figure 7.** Comparison of the estimated hourly PV power generation between simulations performed
 435 using GSEE with input data consisting of either only GHI or both GHI and DHI under varying aerosol
 436 conditions: (top) fixed panels (bottom) 2-axis tracking systems

437 The PV systems considered in this study have a nominal capacity of 1 kWp. The PV model applies a
 438 default system loss factor of 10%. This effectively limits the maximum achievable power output to
 439 approximately 90% of the nominal capacity (i.e., around 900 W/kWp). This effect becomes apparent
 440 at the Izaña site due to its low latitude combined with its specific geographical and atmospheric
 441 conditions, which lead to high irradiance levels. As a result, the simulated PV output in some cases
 442 appears capped around 900 Wh/kWp per hour when only GHI is used.

443 Additionally, Tables 4 and 5 present the validation results for Carpentras and Tamanrasset, selected
 444 as representative locations that encompass a wide variety of sky conditions. Validation results for
 445 the remaining stations are available in the supplement (Tables S1-S3). All the evaluation metrics
 446 correspond to simulations of PV panels with c-Si technology.

447 **Table 4.** Evaluation metrics for GSEE performance within hourly intervals in Carpentras, comparing
 448 simulations with diffuse fraction from measurements and from the BRL model

STATION: Carpentras		fixed panels			2-axis tracking		
		RMSE (Wh/kWp/hour)	MAE (Wh/kWp/hour)	rMBE (%)	RMSE (Wh/kWp/hour)	MAE (Wh/kWp/hour)	rMBE (%)
All-Sky scenes		12.6	6.6	0.8	20.8	12.5	1.2
All-Sky scenes (cloudiness)	cloud-free	9.2	4.6	0.4	14.8	8.7	0.5
	partly cloudy	19.5	12.5	2.3	32.5	23.9	3.8
	cloudy (overcast)	5.8	3.0	2.0	10.5	6.1	4.6
Cloudless- Sky scenes (aerosol load)	low	4.7	3.4	-0.4	9.5	7.5	-0.8
	moderate	4.3	2.2	0.1	7.8	4.7	0.0
	high	6.4	4.0	0.6	11.0	7.8	0.9
	very high	14.9	10.2	1.6	22.7	17.2	2.6

449

450 **Table 5.** Evaluation metrics for GSEE performance within hourly intervals in Tamanrasset,
 451 comparing simulations with diffuse fraction from measurements and from the BRL model.

STATION: Tamanrasset		fixed panels			2-axis tracking		
		RMSE (Wh/kWp/hour)	MAE (Wh/kWp/hour)	rMBE (%)	RMSE (Wh/kWp/hour)	MAE (Wh/kWp/hour)	rMBE (%)
All-Sky scenes		13.6	9.3	1.0	40.4	27.8	3.8

All-Sky scenes (cloudiness)	cloud-free	11.5	8.0	0.8	35.3	23.4	2.9
	partly cloudy	20.1	15.0	2.0	56.1	45.7	8.1
	cloudy (overcast)	8.4	5.2	-0.1	45.3	30.1	11.2
Cloudless-Sky scenes (aerosol load)	low	3.2	2.0	0.2	6.6	4.0	0.3
	moderate	5.4	4.6	0.6	13.0	10.5	1.2
	high	12.5	11.7	1.6	30.1	27.4	3.4
	very high	18.0	16.2	1.9	57.0	49.2	6.8

452

453 Based on the calculated statistical indices, the Root Mean Square Error (RMSE) values for fixed
 454 panels range from 4.7 Wh/kWp/hour (clear sky) to 19.5 Wh/kWp/hour (partly cloudy) in Carpentras,
 455 and from 3.2 to 20.1 Wh/kWp/hour in Tamanrasset. Under very high aerosol loading, RMSE reaches
 456 14.9 and 18.0 Wh/kWp/hour, respectively. For 2-axis tracking systems, RMSE values vary
 457 significantly, ranging from 9.5 to 32.5 Wh/kWp/hour in Carpentras and from 6.6 to 56.1 Wh/kWp/hour
 458 in Tamanrasset, with peaks of 22.7 and 57.0 Wh/kWp/hour under very high aerosol loading
 459 conditions. Similarly, the Mean Absolut Error (MAE) values are generally lower for fixed panels (3.4-
 460 12.5 Wh/kWp//hour in Carpentras, 2.0-15.0 in Tamanrasset) and substantially higher for 2-axis
 461 tracking (7.5-23.9 and 4.0-45.7 Wh/kWp/hour, respectively). Notably in Tamanrasset, MAE values
 462 under very high aerosol loading exceed those observed under partly cloudy conditions, with values
 463 increasing from 15.0 to 16.2 Wh/kWp/hour for fixed panels and from 45.7 to 49.2 Wh/kWp/hour for
 464 2-axis tracking systems. Regarding the relative mean bias (rMBE), this remains mostly within $\pm 4.6\%$
 465 for fixed panels but can reach up to 11.2% for 2-axis tracking, particularly in aerosol-laden
 466 conditions.

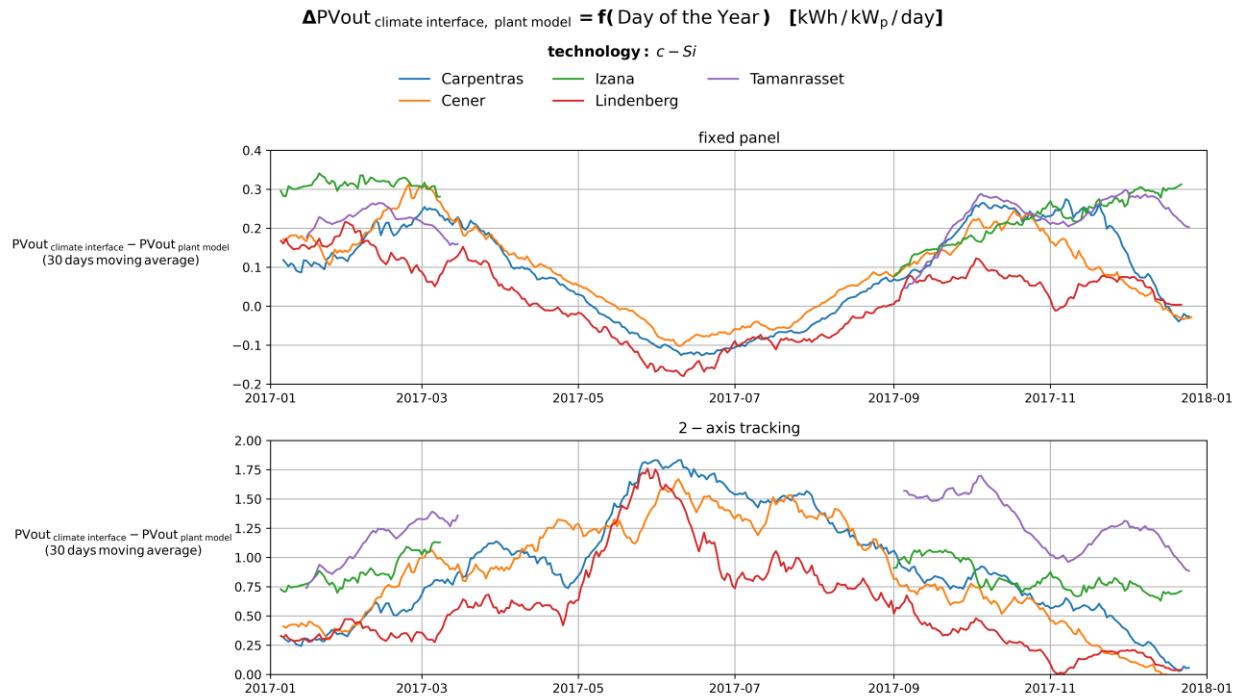
467

468 3.4 Estimating total daily PV power output using the Climate Interface

469 Validation of the estimated daily energy production using the Climate Interface is achieved by
 470 comparing the estimates with the results obtained from the direct summation of the hourly
 471 simulations with input both GHI and DHI.

472 The Climate Interface generates the hourly profile of GHI for each day as a sinusoidal function. Then,
 473 the BRL is applied to the hourly time-series, and the hourly power generation is computed. Finally,
 474 these values are summed up to provide an estimate of the total daily output power. As shown in Fig.
 475 8, which illustrates the differences between the Climate Interface estimates and the sums of the
 476 hourly simulations, this approach introduces a variability throughout the year. Furthermore, Figure

477 S6 in the supplement presents the percentage differences between the two approaches, using the
478 latter as the reference.



479

480 **Figure 8.** Time-series of the differences between the daily PV output estimated using the climate
481 interface and the corresponding daily sums from hourly simulations.

482

483 The time-series represent the centered 30-day moving average. To ensure that the values are
484 representative of the reference period, we have applied all conditions requiring at least 20 days of
485 available data within each 30-days interval. In Tamanrasset and Izaña, especially during the summer
486 months, there are significant data gaps on several days, often occurring around solar noon.

487 More precisely, from Fig. 8, we observe that within the modelling of PV plants with fixed panels, there
488 is a tendency to overestimate in winter, with deviations of approximately 0.3 kWh/kWp/day, and to
489 slightly underestimate in summer, where deviations are around 0.1 kWh/kWp/day. In contrast, for 2-
490 axis solar tracking systems, the resulting deviations are significantly larger, with a general tendency
491 toward overestimation that peaks during summer, reaching approximately 1.75 kWh/kWp/day. The
492 percentage differences span from -10 to 20 % for fixed panels and from -5 to 35 % for 2-axis tracking
493 systems.

494 The variability in the percentage difference between the daily PV output estimated using the climate
495 interface and the corresponding daily sums is mainly a function of the minimum SZA, while
496 especially in the case of modeling for 2-axes tracking systems, the variation is also influenced by
497 aerosol loading, with differences tending to increase as aerosol load rises (Figures S4 and S5 in the
498 supplement).

499 Additional validation results are provided in the supplement (Tables S4-S8). Indicatively, for
500 Carpentras and Tamanrasset, representative results are discussed below. For fixed panels, RMSE is
501 minimized at 0.18 kWh/kWp/day under very-low aerosol conditions, compared to the overall 0.22
502 kWh/kWp/day for Carpentras. In Tamanrasset, the lowest RMSE is observed at 0.15 kWh/kWp/day
503 under very low aerosol conditions, while the overall reaches 0.24. In the case of 2-axis tracking, a
504 significant increase is observed from low-aerosol to aerosol-laden conditions, ranging from 0.82 to
505 1.28 kWh/kWp/day in Carpentras and from 0.66 to 1.37 in Tamanrasset. Similar widening trends are
506 also evident in the MAE values across different aerosol loading conditions. The computed statistical
507 indices confirm that the differences are minimized under sunny and nearly aerosol-free sky
508 conditions. Comparing the performance on low-aerosol days to that on aerosol-laden, we conclude
509 that, particularly in the case of modelling 2-axis tracking systems, errors increase significantly. In
510 Tamanrasset, in particular, the errors are more than double.

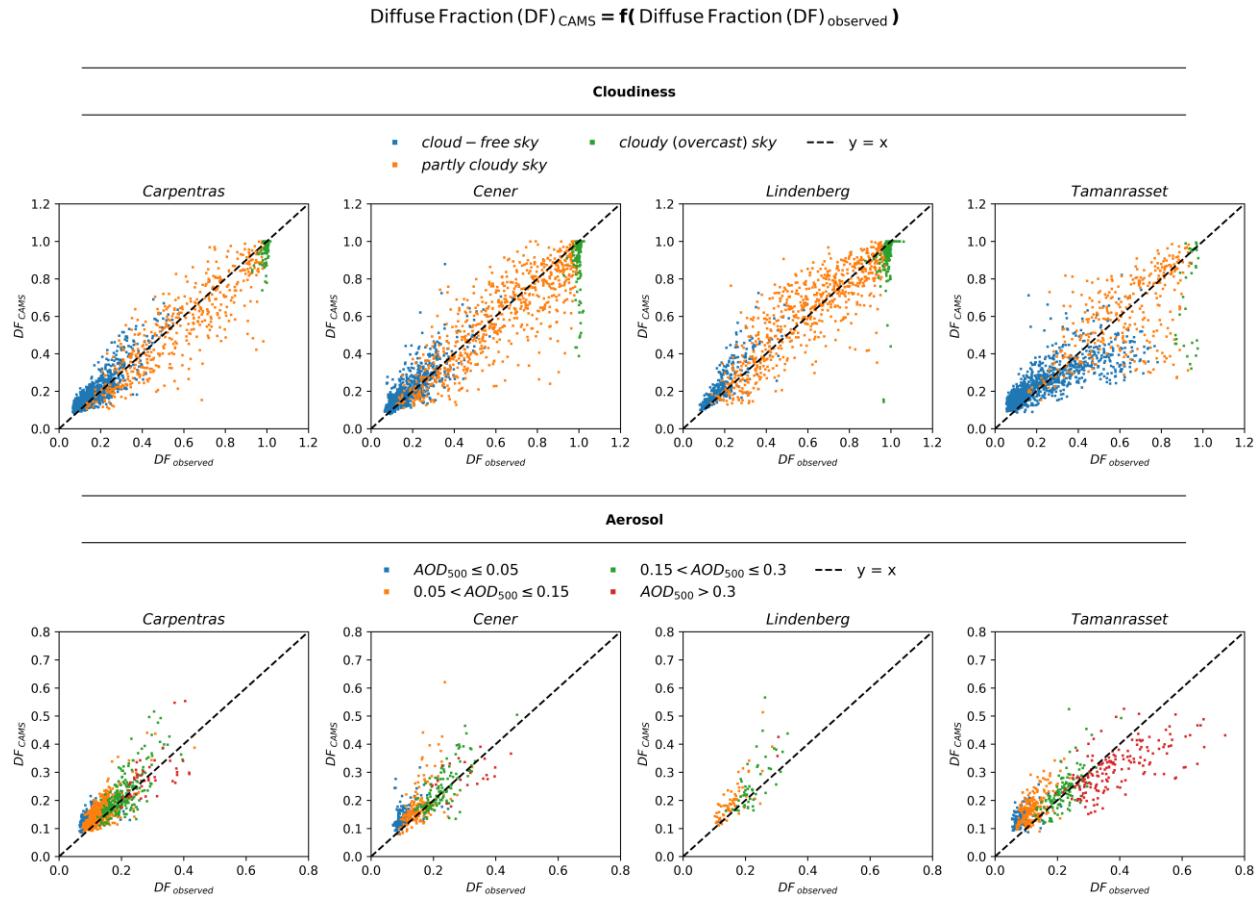
511 3.5 Evaluation of the reliability of using the CAMS solar radiation time-series product in modelling
512 PV power potential

513 The aim of this section is to inspect the reliability of using the CAMS solar radiation time-series
514 product in modelling the PV power potential adapted to a certain location. A review of the existing
515 literature indicates a lack of studies directly examining the accuracy of using CAMS data for
516 assessing PV power potential. This is addressed by comparing the output power obtained from using
517 CAMS solar radiation data with that calculated using ground-based measurements. The analysis
518 focuses on the capability of CAMS to provide accurate estimates of both GHI as well as its individual
519 components.

520 In this section, we have excluded Izaña, because, due to its high altitude – as indicated through a
521 personal communication with Yves-Marie Saint-Drenan (2025) – comparable results would require
522 adjusting the measurements to the elevation of the stations, which is a complicated process and
523 beyond the scope of this study.

524 The CAMS-based diffuse fraction, compared to the observed, is presented in Figure 9 under different
 525 prevailing conditions. We observe that the calculation of the diffuse component is subject to
 526 significant uncertainty. Cloudiness is the primary uncertainty source, particularly under partly cloudy
 527 conditions. Additionally, notable discrepancies related to aerosols emerge only in cases of very high
 528 aerosol loading.

529



530

531 **Figure 9.** Comparison of the CAMS-based diffuse fraction estimated using BRL with the actual one
 532 under diverse atmospheric conditions

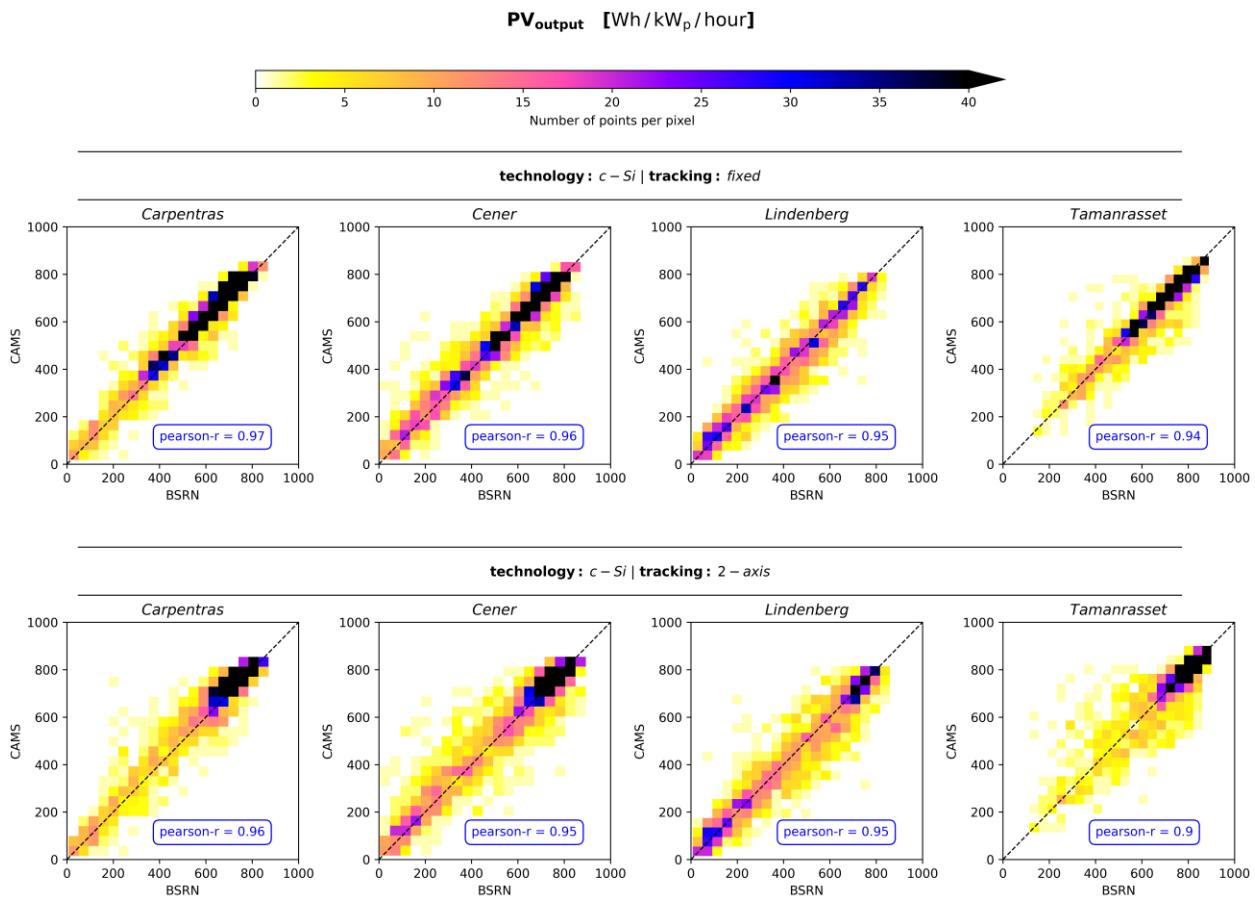
533

534 In Fig. 10 we provide density scatter plots comparing the CAMS-based PV output power with that
 535 computed from the ground-based BSRN data, aiming to illustrate how the uncertainty in the diffuse
 536 component estimates propagate to the calculation of power generation. Notably, there is a much
 537 greater dispersion from the y=x line in the case of simulating PV plants with 2-axis tracking system,

538 compared to that within the modelling of fixed panels. This outcome is attributed to the increased
539 sensitivity of the 2-axis tracking systems to the partitioning of global irradiance into its components.
540 Nevertheless, correlation coefficients are in all cases better than 0.9.

541 Additional evaluation metrics are provided in the supplement (Tables S9-S12). Indicatively, we
542 observe that under cloudless conditions, for fixed panels, RMSE ranges between 25.0 to 42.3
543 Wh/kWp/hour in Carpentras and 16.6 and 31.0 Wh/kWp/hour in Tamanrasset, with variations linked
544 to aerosol loading. Similarly, MAE ranges from 20.0 to 36.9 Wh/kWp/hour in Carpentras and 11.9 to
545 22.9 Wh/kWp/hour in Tamanrasset. For 2-axis systems, RMSE and MAE follow similar trend, ranging
546 from 28.8 to 49.9 Wh/kWp/hour and 22.3 to 44.1 Wh/kWp/hour, respectively, in Carpentras, and from
547 20.8 to 48.0 Wh/kWp/hour and 15.3 to 35.5 Wh/kWp/hour, respectively, in Tamanrasset. Conversely,
548 under cloudy conditions the errors are significantly increasing. In Carpentras, as well as in Cener,
549 and Lindenberg (according to the corresponding tables in the supplement) the errors peak under
550 partly cloudy conditions, with RMSE reaching up to 94.2 Wh/kWp/hour in Carpentras. However, in
551 Tamanrasset, the highest errors occur under overcast conditions, where RMSE and MAE for 2-axis
552 solar tracking systems reach 210.7 and 151.6 Wh/kWp/hour, respectively. This exception can be
553 interpreted through Figure 15, which illustrates that in the rare overcast scenes in Tamanrasset,
554 CAMS occasionally reports low diffuse fraction values instead of values close to 1, suggesting that
555 CAMS did not accurately represent cloudiness in these cases.

556



557

558 **Figure 10.** Overview of the reliability of the CAMS-based PV power simulations

559

560 **4. Conclusions**

561 The optimal approach to include solar radiation information to PV power models such as GSEE is to
 562 use actual in-situ measurements of global and diffuse solar irradiance. Since measurements of the
 563 diffuse component are rarely available, it is common to use measurements of the GHI (if available)
 564 and retrieve the diffuse component using a model such as BRL. In the absence of in-situ
 565 measurements, other options include the use of datasets such as CAMS or even a radiative transfer
 566 model, provided that atmospheric inputs such as clearness index, aerosol optical depth (AOD), and
 567 other aerosol properties are available. This study evaluated these options and their implications for
 568 PV modelling accuracy.

569 The results highlighted the importance of having precise information for the distribution of solar
 570 irradiance among its components in PV power modelling. The implementation of the BRL diffuse

571 fraction within GSEE serves as a practical, and under certain conditions, reliable solution to the
572 absence of detailed information for each component separately. Moreover, the integrated Climate
573 Data Interface submodule offers valuable prospects for investigating fluctuations in the solar PV
574 power generation across various timescales. In this context, the use of BRL has a key contribution
575 alongside the other computational procedures in processing climate datasets. Previous studies on
576 PV power modelling approaches have not examined their reliability under diverse atmospheric
577 conditions, including the effects associated with cloudiness, aerosol loading, as well as aerosol
578 optical properties.

579 The evaluation of the BRL's performance revealed a dependency of its reliability on the prevailing sky
580 conditions. BRL has excellent accuracy under totally clear sky scenes and still performs well for
581 cloudless scenes with moderate aerosol loading. In general, its accuracy is inversely proportional to
582 the complexity of the cloud scene. However, the model systematically underestimates the diffuse
583 fraction under high-loading conditions, such as during dust events. The discrepancies arising from
584 diffuse fraction estimation propagate to PV power generation and become particularly pronounced
585 in the modelling of 2-axis tracking systems. Indicatively, MAE under cloud-free scenes with moderate
586 aerosol loading, ranges between 2.2 to 6.6 Wh/kWp/hour for fixed panels and 4.7 to 15.0
587 Wh/kWp/hour for 2-axis tracking systems. Under partly cloudy conditions, where the cloud scene is
588 more complex, the MAE increases substantially, ranging from 12.4 to 25.8 Wh/kWp/hour for fixed
589 panels and from 23.5 to 55.1 Wh/kWp/hour for 2-axis tracking systems. Moreover, during intense dust
590 events, MAE can reach up to 49.2 Wh/kWp/hour in Tamanrasset, which is comparable to that
591 computed under partly cloudy conditions. Overall, the rMBE remains within the $\pm 5\%$, with the
592 exception of a limited cases under overcast conditions. The same analysis applied to CdTe panels
593 yielded similar results, with minor differences.

594 Aiming to provide an indicative assessment of the financial impacts of the effect of desert dust
595 aerosols, we assume that the statistical indices calculated for Tamanrasset are representative of a
596 large-scale solar farm located in the Sahara region, with 500 MW installed PV capacity and systems
597 equipped with 2-axis solar tracking system. For this hypothetical solar farm, according to the value
598 of the Mean Absolute Error (MAE) on Table 4 for very high aerosol loading, we estimate that the
599 produced energy is $0.0492 [\text{kWh}/\text{kWp}/\text{hour}] \times 500 \times 10^3 [\text{kWp}] = 24600 [\text{kWh}/\text{hour}]$
600 *supposing 12 sunlight hours per day* $\rightarrow \sim 295200 [\text{kWh}/\text{day}]$ less than the expected from the PV power
601 simulations. According to the global average auction prices for selling produced energy back to the

602 grid in 2021 (IRENA, n.d.), the overestimations are equivalent to a financial loss of
603 $0.039 \text{ [USD/kWh]} \times 295200 \text{ [kWh/day]} \approx 11,500 \text{ USD/day}$. Therefore, site assessments that do
604 not correctly account for the distribution of surface solar irradiance in the sky under desert dust
605 aerosol conditions may overestimate financial performance and the annual financial deficit could be
606 accumulated to hundreds of thousands of US dollars per year.

607 Comparing the range of computed errors, we observe that the errors arising from employing CAMS
608 rather than using ground-based measurements, even when the diffuse fraction is not provided, are
609 higher across the overwhelming majority of the considered sky conditions. More specifically,
610 regarding the overall performance, MAE when using CAMS ranges between 33.7 and 46.1
611 Wh/kWp/hour, while with ground-based GHI measurements, MAE remains below 10 Wh/kWp/hour
612 within the modelling of systems with fixed panels and can reach up to 27.8 Wh/kWp/hour within the
613 modelling of 2-axis tracking systems. This outcome highlights the value of ground-based
614 measurements.

615 To sum up, achieving the highest quality PV power simulations necessitates high-quality, concurrent
616 measurements of solar irradiance components. In absence of this, the submodules included in the
617 GSEE package enable reliable simulations under the vast majority of prevailing sky conditions. CAMS
618 serves as a valuable data source for PV power modelling, but it cannot fully replace the precision and
619 reliability of using ground-based measurements. The integration of aerosol correction within the BRL
620 model opens new possibilities for further improvements in the modelling of solar energy systems. A
621 more comprehensive assessment would require measured PV output data; however, acquiring
622 simultaneous direct and diffuse irradiance measurements at the same location as the solar farms
623 remains challenging.

624

625 **Data availability**

626 The BSRN data are freely available on the BSRN web-page (<https://bsrn.awi.de/>). The AERONET
627 version 3 products are freely available from the AERONET website (<https://aeronet.gsfc.nasa.gov/>).
628 The CAMS radiation time-series are available from the Atmosphere Data Store
629 (<https://ads.atmosphere.copernicus.eu>). The rest of the data used in this paper are available upon
630 request from the authors.

631 **Author Contributions**

632 Conceptualization: NP and IF; Data curation: NP and KP; Formal analysis: NP; Funding acquisition:
633 CZ; Investigation: NP; Methodology: NP, IF, SK, AK and AG; Project administration: CZ; Resources: SP,
634 KP and LD; Software: NP; Supervision: IF; Validation: NP, IF and SP; Visualization: NP; Writing –
635 original draft: NP; Writing – review & editing: all authors

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650 **References**

651 Anderson, G., Clough, S., Kneizys, F., Chetwynd, J., & Shettle, E. (1986). *AFGL atmospheric*
652 *constituent profiles (0-120 km)* (Tech. Rep. AFGL-TR-86-0110). Air Force Geophysics Laboratory,
653 Hanscom Air Force Base.

654 Anderson, K. S., Hansen, C. W., Holmgren, W. F., Jensen, A. R., Mikofski, M. A., & Driesse, A.
655 (2023). *pvlb python: 2023 project update. Journal of Open Source Software*, 8(92), Article 5994.
656 <https://doi.org/10.21105/joss.05994>

657 Ångström, A. (1929). On the atmospheric transmission of sun radiation and on dust in the
658 air. *Geografiska Annaler*, 11(2), 156–166. <https://doi.org/10.1080/20014422.1929.11880498>

659 Barreto, Á., García, R. D., Guirado-Fuentes, C., Cuevas, E., Almansa, A. F., Milford, C., Toledano,
660 C., Expósito, F. J., Díaz, J. P., & León-Luis, S. F. (2022). Aerosol characterisation in the
661 subtropical eastern North Atlantic region using long-term AERONET
662 measurements. *Atmospheric Chemistry and Physics*, 22(17), 11105–11124.
663 <https://doi.org/10.5194/acp-22-11105-2022>

664 Blaga, R., Mares, O., Paulescu, E., Boata, R., Sabadus, A., Hategan, S.-M., Calinou, D., Stefu,
665 N., & Paulescu, M. (2024). Diffuse fraction as a tool for exploring the sensitivity of parametric
666 clear-sky models to changing aerosol conditions. *Solar Energy (Phoenix, Ariz.)*, 277(112731),
667 112731. <https://doi.org/10.1016/j.solener.2024.112731>

668 Blanc, P., Remund, J., & Vallance, L. (2017). Short-term solar power forecasting based on
669 satellite images. In *Renewable Energy Forecasting* (pp. 179–198). Elsevier.

670 Boland, J. W., Scott, L., & Luther, M. (2001). *Modelling the diffuse fraction of global solar
671 radiation on a horizontal surface*. *Environmetrics*, 12(2), 103–116. [https://doi.org/10.1002/1099-095X\(200103\)12:2<103::AID-ENV447>3.0.CO;2-2](https://doi.org/10.1002/1099-
672 095X(200103)12:2<103::AID-ENV447>3.0.CO;2-2)

673 Buras, R., Dowling, T., & Emde, C. (2011). New secondary-scattering correction in DISORT with
674 increased efficiency for forward scattering. *Journal of Quantitative Spectroscopy & Radiative
675 Transfer*, 112(12), 2028–2034. <https://doi.org/10.1016/j.jqsrt.2011.03.019>

676 Cañadillas-Ramallo, D., Moutaoikil, A., Shephard, L. E., & Guerrero-Lemus, R. (2022). The
677 influence of extreme dust events in the current and future 100% renewable power scenarios in
678 Tenerife. *Renewable Energy*, 184, 948–959. <https://doi.org/10.1016/j.renene.2021.12.013>

679 Copernicus Atmosphere Monitoring Service. (2020). *CAMS solar radiation time-series*.
680 Copernicus Atmosphere Monitoring Service (CAMS) Atmosphere Data Store.
681 <https://doi.org/10.24381/5cab0912>

682 Cuevas, E., Romero-Campos, P. M., Kouremeti, N., Kazadzis, S., Räisänen, P., García, R. D.,
683 Barreto, A., Guirado-Fuentes, C., Ramos, R., Toledano, C., Almansa, F., & Gröbner, J. (2019).
684 Aerosol optical depth comparison between GAW-PFR and AERONET-Cimel radiometers from
685 long-term (2005–2015) 1 min synchronous measurements. *Atmospheric Measurement
686 Techniques*, 12(8), 4309–4337. <https://doi.org/10.5194/amt-12-4309-2019>

687 Driemel, A., Augustine, J., Behrens, K., Colle, S., Cox, C., Cuevas-Agulló, E., Denn, F. M., Duprat,
688 T., Fukuda, M., Grobe, H., Haeffelin, M., Hodges, G., Hyett, N., Ijima, O., Kallis, A., Knap, W.,
689 Kustov, V., Long, C. N., Longenecker, D., ... König-Langlo, G. (2018). Baseline Surface Radiation
690 Network (BSRN): structure and data description (1992–2017). *Earth System Science
691 Data*, 10(3), 1491–1501. <https://doi.org/10.5194/essd-10-1491-2018>

692 Dubey, S., Sarvaiya, J. N., & Seshadri, B. (2013). Temperature dependent photovoltaic (PV)
693 efficiency and its effect on PV production in the world – A review. *Energy Procedia*, 33, 311–321.
694 <https://doi.org/10.1016/j.egypro.2013.05.072>

695 Dubovik, O., Holben, B., Eck, T. F., Smirnov, A., Kaufman, Y. J., King, M. D., Tanré, D., & Slutsker, I.
696 (2002). Variability of absorption and optical properties of key aerosol types observed in
697 worldwide locations. *Journal of the Atmospheric Sciences*, 59(3), 590–608.
698 [https://doi.org/10.1175/1520-0469\(2002\)059<0590:voaaop>2.0.co;2](https://doi.org/10.1175/1520-0469(2002)059<0590:voaaop>2.0.co;2)

699 Dubovik, O., & King, M. D. (2000). A flexible inversion algorithm for retrieval of aerosol optical
700 properties from Sun and sky radiance measurements. *Journal of Geophysical
701 Research*, 105(D16), 20673–20696. <https://doi.org/10.1029/2000jd900282>

702 Edensofer, O., Pichs-Madruga, R., Sokona, Y., Seyboth, K., Kadner, S., Zwickel, T., Eickemeier, P.,
703 Hansen, G., Schlomer, S., & Von Stechow, C. (Eds.). (2011). *Renewable energy sources and
704 climate change mitigation: Special report of the intergovernmental panel on climate change*.
705 Cambridge University Press.

706 Emde, C., Buras-Schnell, R., Kylling, A., Mayer, B., Gasteiger, J., Hamann, U., Kylling, J., Richter,
707 B., Pause, C., Dowling, T., & Bugliaro, L. (2016). The libRadtran software package for radiative
708 transfer calculations (version 2.0.1). *Geoscientific Model Development*, 9(5), 1647–1672.
709 <https://doi.org/10.5194/gmd-9-1647-2016>

710 Faid, A., Smara, Y., Caselles, V., & Khireddine, A. (2012). Evaluation of the Saharan aerosol
711 impact on solar radiation over the Tamanrasset area, Algeria. *International Journal of Advanced
712 Research in Engineering and Technology*, 3(1), 24–32.

713 Fountoulakis, I., Kosmopoulos, P., Papachristopoulou, K., Raptis, I.-P., Mamouri, R.-E., Nisantzi,
714 A., Gkikas, A., Witthuhn, J., Bley, S., Moustaka, A., Buehl, J., Seifert, P., Hadjimitsis, D. G.,
715 Kontoes, C., & Kazadzis, S. (2021). Effects of aerosols and clouds on the levels of surface solar

716 radiation and solar energy in Cyprus. *Remote Sensing*, 13(12), 2319.

717 <https://doi.org/10.3390/rs13122319>

718 Fountoulakis, I., Papachristopoulou, K., Proestakis, E., Amiridis, V., Kontoes, C., & Kazadzis, S.

719 (2022). Effect of aerosol vertical distribution on the modeling of solar radiation. *Remote*

720 *Sensing*, 14(5), 1143. <https://doi.org/10.3390/rs14051143>

721 Giles, D. M., Sinyuk, A., Sorokin, M. G., Schafer, J. S., Smirnov, A., Slutsker, I., Eck, T. F., Holben,

722 B. N., Lewis, J. R., Campbell, J. R., Welton, E. J., Korkin, S. V., & Lyapustin, A. I. (2019).

723 Advancements in the Aerosol Robotic Network (AERONET) Version 3 database – automated

724 near-real-time quality control algorithm with improved cloud screening for Sun photometer

725 aerosol optical depth (AOD) measurements. *Atmospheric Measurement Techniques*, 12(1),

726 169–209. <https://doi.org/10.5194/amt-12-169-2019>

727 Holben, B. N., Eck, T. F., Slutsker, I., Tanré, D., Buis, J. P., Setzer, A., Vermote, E., Reagan, J. A.,

728 Kaufman, Y. J., Nakajima, T., Lavenu, F., Jankowiak, I., & Smirnov, A. (1998). AERONET—A

729 federated instrument network and data archive for aerosol characterization. *Remote Sensing of*

730 *Environment*, 66(1), 1–16. [https://doi.org/10.1016/s0034-4257\(98\)00031-5](https://doi.org/10.1016/s0034-4257(98)00031-5)

731 Hou, X., Wild, M., Folini, D., Kazadzis, S., & Wohland, J. (2021). Climate change impacts on solar

732 power generation and its spatial variability in Europe based on CMIP6. *Earth System*

733 *Dynamics*, 12(4), 1099–1113. <https://doi.org/10.5194/esd-12-1099-2021>

734 Huld, T., Gottschalg, R., Beyer, H. G., & Topič, M. (2010). Mapping the performance of PV

735 modules, effects of module type and data averaging. *Solar Energy (Phoenix, Ariz.)*, 84(2), 324–

736 338. <https://doi.org/10.1016/j.solener.2009.12.002>

737 Intergovernmental Panel on Climate Change (IPCC). (2023). *Climate change 2022 – impacts,*

738 *adaptation and vulnerability: Working group II contribution to the sixth assessment report of the*

739 *intergovernmental panel on climate change*. Cambridge University Press.

740 <https://doi.org/10.1017/9781009325844>

741 Jacovides, C. P., Tymvios, F. S., Assimakopoulos, V. D., & Kaltsounides, N. A. (2006).

742 Comparative study of various correlations in estimating hourly diffuse fraction of global solar

743 radiation. *Renewable Energy*, 31(15), 2492–2504. <https://doi.org/10.1016/j.renene.2005.11.009>

744 Kakran, S., Rathore, J. S., Sidhu, A., & Kumar, A. (2024). Solar energy advances and CO₂
745 emissions: A comparative review of leading nations' path to sustainable future. *Journal of*
746 *Cleaner Production*, 475(143598), 143598. <https://doi.org/10.1016/j.jclepro.2024.143598>

747 Kato, S., Ackerman, T. P., Mather, J. H., & Clothiaux, E. E. (1999). The k-distribution method and
748 correlated-k approximation for a shortwave radiative transfer model. *Journal of Quantitative*
749 *Spectroscopy & Radiative Transfer*, 62(1), 109–121. [https://doi.org/10.1016/s0022-4073\(98\)00075-2](https://doi.org/10.1016/s0022-4073(98)00075-2)

751 Kazantzidis, A., Tzoumanikas, P., Blanc, P., Massip, P., Wilbert, S., & Ramirez-Santigosa, L.
752 (2017). Short-term forecasting based on all-sky cameras. In *Renewable Energy Forecasting* (pp.
753 153–178). Elsevier.

754 Kosmopoulos, P., Kazadzis, S., El-Askary, H., Taylor, M., Gkikas, A., Proestakis, E., Kontoes, C., &
755 El-Khayat, M. (2018). Earth-Observation-based estimation and forecasting of particulate matter
756 impact on solar energy in Egypt. *Remote Sensing*, 10(12), 1870.
757 <https://doi.org/10.3390/rs10121870>

758 Kouklaki, D., Kazadzis, S., Raptis, I.-P., Papachristopoulou, K., Fountoulakis, I., & Eleftheratos,
759 K. (2023). Photovoltaic spectral responsivity and efficiency under different aerosol
760 conditions. *Energies*, 16(18), 6644. <https://doi.org/10.3390/en16186644>

761

762

763 Lauret, P., Boland, J., & Ridley, B. (2013). Bayesian statistical analysis applied to solar radiation
764 modelling. *Renewable Energy*, 49, 124–127. <https://doi.org/10.1016/j.renene.2012.01.049>

765 Liu, B. Y. H., & Jordan, R. C. (1960). The interrelationship and characteristic distribution of
766 direct, diffuse and total solar radiation. *Solar Energy (Phoenix, Ariz.)*, 4(3), 1–19.
767 [https://doi.org/10.1016/0038-092x\(60\)90062-1](https://doi.org/10.1016/0038-092x(60)90062-1)

768 Logothetis, S.-A., Salamalikis, V., & Kazantzidis, A. (2020). Aerosol classification in Europe,
769 Middle East, North Africa and Arabian Peninsula based on AERONET Version 3. *Atmospheric*
770 *Research*, 239(104893), 104893. <https://doi.org/10.1016/j.atmosres.2020.104893>

771 Long, C., & Dutton, E. (2010). *BSRN Global Network recommended QC tests, V2.x*.
772 https://epic.awi.de/id/eprint/30083/1/BSRN_recommended_QC_tests_V2.pdf

773 Mayer, B., & Kylling, A. (2005). Technical note: The libRadtran software package for radiative
774 transfer calculations - description and examples of use. *Atmospheric Chemistry and*
775 *Physics*, 5(7), 1855–1877. <https://doi.org/10.5194/acp-5-1855-2005>

776 McMahan, A. C., Grover, C. N., & Vignola, F. E. (2013). Evaluation of resource risk in solar-
777 project financing. In *Solar Energy Forecasting and Resource Assessment* (pp. 81–95). Elsevier.

778 Owusu, P. A., & Asumadu-Sarkodie, S. (2016). A review of renewable energy sources,
779 sustainability issues and climate change mitigation. *Cogent Engineering*, 3(1), 1167990.
780 <https://doi.org/10.1080/23311916.2016.1167990>

781 Papachristopoulou, K., Fountoulakis, I., Bais, A. F., Psiloglou, B. E., Papadimitriou, N., Raptis, I.-
782 P., Kazantzidis, A., Kontoes, C., Hatzaki, M., & Kazadzis, S. (2024). Effects of clouds and
783 aerosols on downwelling surface solar irradiance nowcasting and short-term
784 forecasting. *Atmospheric Measurement Techniques*, 17(7), 1851–1877.
785 <https://doi.org/10.5194/amt-17-1851-2024>

786 Papachristopoulou, K., Fountoulakis, I., Gkikas, A., Kosmopoulos, P. G., Nastos, P. T., Hatzaki,
787 M., & Kazadzis, S. (2022). 15-year analysis of direct effects of total and dust aerosols in solar
788 radiation/energy over the Mediterranean Basin. *Remote Sensing*, 14(7), 1535.
789 <https://doi.org/10.3390/rs14071535>

790 Paulescu, E., & Blaga, R. (2019). A simple and reliable empirical model with two predictors for
791 estimating 1-minute diffuse fraction. *Solar Energy (Phoenix, Ariz.)*, 180, 75–84.
792 <https://doi.org/10.1016/j.solener.2019.01.029>

793 Pedro, H. T. C., Inman, R. H., & Coimbra, C. F. M. (2017). Mathematical methods for optimized
794 solar forecasting. In *Renewable Energy Forecasting* (pp. 111–152). Elsevier.

795 Pfenninger, S., & Staffell, I. (2016). Long-term patterns of European PV output using 30 years of
796 validated hourly reanalysis and satellite data. *Energy (Oxford, England)*, 114, 1251–1265.
797 <https://doi.org/10.1016/j.energy.2016.08.060>

798 Raptis, I.-P., Kazadzis, S., Fountoulakis, I., Papachristopoulou, K., Kouklaki, D., Psiloglou, B. E.,
799 Kazantzidis, A., Benetatos, C., Papadimitriou, N., & Eleftheratos, K. (2023). Evaluation of the
800 solar energy nowcasting system (SENSE) during a 12-months intensive measurement campaign
801 in Athens, Greece. *Energies*, 16(14), 5361. <https://doi.org/10.3390/en16145361>

802

803 Qu, Z., Oumbe, A., Blanc, P., Espinar, B., Gesell, G., Gschwind, B., Klüser, L., Lefèvre, M.,
804 Saboret, L., Schroedter-Homscheidt, M., & Wald, L. (2017). Fast radiative transfer
805 parameterisation for assessing the surface solar irradiance: The Heliosat-4
806 method. *Meteorologische Zeitschrift*, 26(1), 33–57. <https://doi.org/10.1127/metz/2016/0781>

807 Renewables Ninja. (n.d.). *Climate data interface*. GSEE Documentation.
808 <https://gsee.readthedocs.io/en/latest/climatedata-interface/>

809 Ridley, B., Boland, J., & Lauret, P. (2010). Modelling of diffuse solar fraction with multiple
810 predictors. *Renewable Energy*, 35(2), 478–483. <https://doi.org/10.1016/j.renene.2009.07.018>

811 Schroedter-Homscheidt, M., Azam, F., Betcke, J., Hanrieder, N., Lefèvre, M., Saboret, L., &
812 Saint-Drenan, Y. -M. (2022). Surface solar irradiation retrieval from MSG/SEVIRI based on
813 APOLLO Next Generation and HELIOSAT-4 methods. *Meteorologische Zeitschrift*, 31(6), 455–
814 476. <https://doi.org/10.1127/metz/2022/1132>

815 Shettle, E. (1989). Models of aerosols, clouds, and precipitation for atmospheric propagation
816 studies. In *Atmospheric propagation in the UV, visible, IR and mm-region and related system*
817 *aspects* (AGARD Conference Proceedings No. 454). NATO Advisory Group for Aerospace
818 Research and Development.

819 Stoffel, T. (2013). Terms and Definitions. In *Solar Energy Forecasting and Resource*
820 *Assessment* (pp. 1–19). Elsevier.

821 Toledano, C., González, R., Fuertes, D., Cuevas, E., Eck, T. F., Kazadzis, S., Kouremeti, N.,
822 Gröbner, J., Goloub, P., Blarel, L., Román, R., Barreto, Á., Berjón, A., Holben, B. N., & Cachorro,
823 V. E. (2018). Assessment of Sun photometer Langley calibration at the high-elevation sites
824 Mauna Loa and Izaña. *Atmospheric Chemistry and Physics*, 18(19), 14555–14567.
825 <https://doi.org/10.5194/acp-18-14555-2018>

826 WMO. (2021). *Guide to instruments and methods of observation* (WMO-No. 8).
827 https://library.wmo.int/doc_num.php?explnum_id=57838

828 Yang, D. (2019). SolarData package update v1.1: R functions for easy access of Baseline
829 Surface Radiation Network (BSRN). *Solar Energy (Phoenix, Ariz.)*, 188, 970–975.
830 <https://doi.org/10.1016/j.solener.2019.05.068>

831

832

833

834