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SPREADS: From Research to Operational Open-Source Data Assimilation System

Carla Cardinali¹

Giovanni Conti¹

Marcelo Guatura¹

Sami Saarinen²

Luis Gustavo Gonçalves De Gonçalves¹

Jeffrey Anderson³

Kevin Raeder³

1- CMCC-Foundation Euro-Mediterranean Center on Climate Change, Italy

2-HPC-consulting Ltd, Finland

3- NSF National Center for Atmospheric Research (NSF-NCAR), Boulder, Colorado, USA

Corresponding author: Carla Cardinali, carla.cardinali@cmcc.it



Abstract

The Scalable PaRallelised EArth Data Assimilation System (SPREADS) is a next-generation data assimilation system developed at CMCC-Foundation (Euro-Mediterranean Center on Climate Change) to support operational global forecasts. Built upon the Data Assimilation Research Testbed (DART), SPREADS incorporates key advancements such as First Guess at Appropriate Time (FGAT), enhanced observation handling via D4O (Database for Observations), and high-performance parallelisation to significantly improve computational efficiency and scalability. A major focus of SPREADS is the assimilation of a vastly increased number of asynchronous satellite-based radiances, which have been shown to substantially enhance analysis quality. Designed for coupled atmosphere-land-ocean-ice assimilation, SPREADS forms the core of the CMCC Earth SYstem Modelling and Data Assimilation (ESYDA division) operational forecast system. This paper presents the modifications made to DART, evaluates preliminary results, and outlines future developments toward fully coupled data assimilation.

1. Introduction & Motivation

Data Assimilation Research Testbed (DART; The Data Assimilation Research Testbed, Version X.Y.Z, 2021, Boulder, Colorado UCAR/NSF NCAR/CISL/DAReS, <http://doi.org/10.5065/D6WQ0202>), is a widely used data assimilation system in the atmospheric and oceanic sciences. It was developed and is maintained by the Data Assimilation Research Section (DAReS) at the NSF National Center for Atmospheric Research (NCAR) in the United States. DART provides a flexible and modular platform for conducting research on data assimilation algorithms and their applications to numerical weather prediction, climate modelling, and other environmental forecasting systems. It primarily focuses on ensemble-based data assimilation methods, such as the Ensemble Kalman Filter (EnKF) and its variants (Evensen 1994a,b; Evensen 2001; Tippett *et al.*, 2003; Collins 2007). These methods use an ensemble of model state vectors to represent the uncertainty in the system state and assimilate observations to update this ensemble. DART is designed with a modular architecture, allowing users to easily integrate different numerical models,



57 observation types, data assimilation algorithms, and experimental configurations.
58 This flexibility enables researchers to tailor the system to specific research questions
59 and applications. DART incorporates the localisation technique (Hamill *et al.* 2001;
60 Houtekamer and Mitchell 2001; Haugen and Evensen 2002 Otto *et al.* 2004) to account
61 for the fact that observations are often only informative within a limited spatial and
62 temporal range, helping therefore to prevent spurious long-range correlations in the
63 analysis, improving the accuracy of the assimilation. DART also includes methods for
64 adaptively inflating the ensemble spread to account for underestimation or
65 overestimation of forecast error covariance. Adaptive inflation is crucial for
66 maintaining the reliability of the ensemble and preventing filter divergence. The
67 benefit to the scientific community of using DART all over these years is without any
68 doubts: DART has been used by a large world wide young and senior researchers to
69 advance understanding data assimilation methods and observations usage (Noh *et al.*
70 2024; Tang *et al.* 2024; Dietrich *et al.* 2024; Pedatella and Anderson 2022; Fox *et al.*
71 2022; Rackza *et al.* 2021). Over the last 20 years, DART has been continuously
72 developed and improved with input and feedback from the scientific community
73 across the atmospheric, oceanic and land sciences. Researchers have contributed new
74 algorithms, techniques, and methodologies expanding the capabilities of the
75 framework and enabling new users to experiment with state-of-the-art approaches
76 (Grooms and Riedel 2024; Anderson 2023; Dibia 2023). By experimenting with
77 various observation types, processing techniques, and quality control methods, users
78 have contributed to optimising the assimilation of observational data and the
79 computational efficiency of DART. This includes improvements in parallelisation
80 strategies, algorithmic optimisations, and enhancements to reduce memory usage and
81 computational costs. And finally also improvements have been obtained on
82 diagnostics, metrics, and benchmarking datasets to assess the quality and reliability
83 of the produced analyses.

84 In 2021 the Euro-Mediterranean Center on Climate Change (CMCC, Italy)
85 approved a new strategy on longer forecast range predictions (e.g. the seasonal
86 forecast) strongly supporting the use of a proper initialisation of these predictions by
87 a weakly coupled data assimilation system. The CMCC strategy foresaw therefore the
88 development of a weakly coupled atmosphere, land, ocean and cryosphere data
89 assimilation system initialising such predictions. Given the crucial role that the open-



90 source modelling plays in terms of transparency and reproducibility, fostering
91 collaboration and community engagement that encourages knowledge sharing, idea
92 exchange, and collective problem-solving, leading to the development of more robust
93 and comprehensive models, CMCC has engaged in the development of a data
94 assimilation system that will serve as an open-source system for operational use.
95 Therefore starting from the open-source DART, SPREADS has evolved by modifying
96 and implementing new features essential for an operational use of the system
97 (Cardinali *et al.*, 2025). CMCC's development of SPREADS as an open-source data
98 assimilation system for operational use builds upon the strengths of the DART
99 framework while customising and extending it to meet the specific requirements of
100 operational forecasting and decision support. In this paper, the description of the
101 changes adopted towards an operational use of an atmospheric data assimilation
102 system is described and assessed. This paper describes the methodological
103 innovations in SPREADS, evaluates its preliminary performance, and outlines future
104 expansion plans toward a coupled open-source DA system.

105

106 **2. Ensemble Kalman Filter and SPREADS**

107

108 The Ensemble Adjustment Kalman Filter (EAKF) is a data assimilation technique
109 developed by members of the DAREs team (Anderson 2001, 2003; Andersson 2009;
110 Andersson, 2012; Anderson and Collins 2012; Reader et al. 2012) within DART. The
111 EAKF addresses key limitations of the standard EnKF, particularly when dealing with
112 small ensemble sizes or poorly known model and observation error statistics. It
113 employs a least squares method to adjust the ensemble state, ensuring consistency
114 with both model dynamics and observational constraints. The EAKF refines the
115 ensemble mean and spread to better fit incoming observations. Observations and
116 ensemble members are assimilated within localised regions, reducing the impact of
117 spurious long-range correlations and enhancing computational efficiency.

118 When new observations become available, the EAKF assimilation process follows
119 two main steps:

- 120 1. Observation-Space Update (Scalar Update): for each observation, the
121 ensemble members are first updated in the observation space. This involves
122 adjusting the prior observation estimates for each ensemble member based on the



123 observed value, the prior ensemble's variance in observation space, and the
 124 observation error variance. This step ensures that the updated observation
 125 estimates are consistent with the new measurement. This update is often
 126 performed as a series of scalar updates if observations are assimilated
 127 sequentially.

128 2. State-Space Adjustment (Ensemble Member Transformation): following the
 129 observation-space update, each ensemble member's state variables are then
 130 adjusted to reflect the change made in the observation space, ensuring that the
 131 updated state remains consistent with the updated observation and the ensemble's
 132 internal correlations. This adjustment explicitly leverages the cross-
 133 covariances between each state variable and the observed variable (computed
 134 directly from the prior ensemble). The transformation ensures that the ensemble
 135 mean and covariance are updated according to the Kalman filter equations, and
 136 crucially, that the updated ensemble members retain their statistical spread and
 137 do not collapse. This adjustment is applied to each ensemble member
 138 independently, guaranteeing that the analysis ensemble still represents the
 139 posterior uncertainty.

140 By explicitly incorporating the least squares assumption, the EAKF provides a
 141 computationally efficient solution. Under these assumptions, the ensemble filtering
 142 problem reduces to a nonlinear filter applied to a scalar, followed by sequential linear
 143 regressions. While subsets of observations with independent error distributions can
 144 be assimilated in sequence, the sequential nature of the regression step presents a
 145 computational challenge when millions of observations must be processed within a
 146 six-hour window. It is well known that satellite radiance assimilation has significantly
 147 improved the quality of numerical weather prediction (NWP) analyses: incorporating
 148 vast amounts of satellite observations, it leads to better initial conditions for
 149 forecasting. To efficiently handle the assimilation of large volumes of satellite
 150 radiances, SPREADS has introduced several modifications, including the First Guess at
 151 Appropriate Time (FGAT) approach, the implementation of RTTOV (the Radiative
 152 Transfer for TIROS Operational Vertical Sounder, Saunders et al., 2018) for radiance
 153 processing (Kugler et al. 2023), the scan and air mass bias corrections, and code
 154 improvements in the observations treatment.

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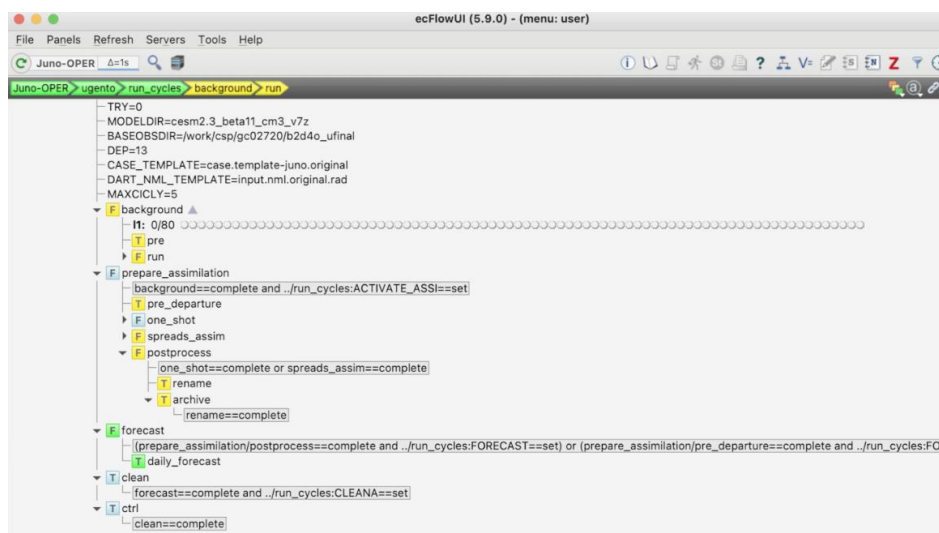


Figure 1: Graphical User Interface EcFlow managing the atmosphere and land component data assimilation and forecast of SPREADS

2.1 Graphical User Interface and Diagnostics

To support the execution of SPREADS, an ensemble data assimilation system processing over one million observations every six hours, a dedicated Graphical User Interface (GUI) was essential. Given the complexity of SPREADS, which involves numerous interdependent modules and programs, a robust and dynamic workflow management tool was required. For this reason, an EcFlow-based GUI (ECMWF EcFlow User Documentation <https://confluence.ecmwf.int/display/ECFLOW>) was developed in parallel with SPREADS. This client/server interface allows for controlled and coordinated execution of all components of the data assimilation suite. The GUI handles task scheduling, monitors job statuses, and responds to events via embedded script commands. It is designed to tolerate hardware or software failures and supports automatic restarts when needed. The GUI manages task dependencies using a trigger-based system, where the execution of one job can depend on the status of others. Job statuses typically include submission, queuing, running, failed, or suspended. These functionalities are enabled by a combination of command-line executables, shared libraries, and a Python-based interface that defines the suite structure and handles communication with the EcFlow server. The server component acts as the scheduler, responding to client requests and managing job execution. While



174 not a queuing system itself, it is capable of submitting jobs to external queueing
175 systems, making it suitable for heterogeneous computing environments. The GUI also
176 provides real-time monitoring and visualisation of the suite's hierarchical node tree,
177 giving users full visibility and control over the operational workflow. Figure 1 shows
178 the EcFlow GUI for SPREADS.

179 Alongside the EcFlow GUI, an interactive graphical diagnostics package based on
180 Streamlit (Streamlit 2024) has been developed to monitor and assess the performance
181 of SPREADS. This tool allows users to explore both model and observation spaces
182 through a range of visualisations. In model space, users can generate geographical
183 maps, cross sections, and mean vertical profiles of key variables. In observation space,
184 the package offers time series and vertical profiles of key statistics, such as biases and
185 standard deviations of the differences between observations and both the prior and
186 posterior fields. Additional diagnostic and dynamical metrics are also available. All the
187 plots presented in this paper were produced using this diagnostic tool.

188

189 2.2 FGAT-approach & Code Modularity

190 One of the key features developed in SPREADS is the FGAT approach, designed to
191 enhance both efficiency and accuracy in the assimilation process. In the First Guess at
192 Appropriate Time (FGAT) approach, the model's background (first-guess) forecast is
193 interpolated to the exact time of each observation, so that the observation-minus-
194 background difference is evaluated at the proper moment before the analysis is
195 performed, yielding a more consistent estimate of the system's current state.

196 In SPREADS, the model's first guess is interpolated to match the time of
197 observational data. To achieve this, the commonly used 6-hour assimilation window
198 is divided into 11 time slots of 30 minutes and 2 additional slots of 15 minutes at the
199 beginning and end of the window. This fine-grained temporal segmentation ensures
200 precise alignment between observations and model output, enabling a more
201 meaningful and accurate comparison.

202 This approach offers several key advantages: it eliminates the need to time-shift
203 either field, which (a) removes interpolation-induced errors and (b) saves
204 considerable CPU and memory, because one forecast integration serves all
205 observation times instead of many separate. Also it enhances consistency by ensuring



206 alignment between the model's initial conditions and observational data, enabling
207 therefore the assimilation of asynchronous observations such as polar orbiting
208 satellite-based observations. Finally, since FGAT can be applied across various data
209 assimilation techniques, it provides flexibility in adapting to different modelling and
210 observational setups.

211 The effectiveness of FGAT was demonstrated during the development of the
212 variational data assimilation system at ECMWF, where the transition from a 6-hour
213 3DVar window to FGAT-3DVar resulted in the largest improvement in assimilation
214 performance (Andersson et al., 1998). The modularity inherent in DART has been
215 further enhanced in SPREADS by refining and structuring the assimilation steps into
216 four distinct modules, each designed to improve efficiency, computational
217 organisation and flexibility. Notably, these modules can be executed independently,
218 allowing for greater adaptability in different assimilation workflows.

219 *Module-0:* Executes the model trajectory using the FGAT approach, while
220 independently handling observation preprocessing.

221 *Module-1:* Performs observation preprocessing, including cross-checking
222 observations, converting them from buffer storage to an SQL query-based database,
223 and conducting screening and blacklisting. Also the scan and air mass bias corrections
224 are performed in here: biases in satellite radiance observations arise due to
225 instrument calibration errors, radiative transfer modeling inaccuracies, and
226 atmospheric variations. The implementation of bias corrections includes adjusting
227 radiances for systematic errors associated with the sensor's viewing angle and using
228 predictors such as atmospheric thickness and surface temperature to adjust radiances
229 for biases linked to atmospheric conditions (Harris and Kelly, 2001, Auligné *et al.* ,
230 2007).

231 *Module-2:* Carries out the nonlinear spatial interpolation of model values from all
232 ensemble members to the observation locations.

233 *Module-3:* Executes the two-step sequential regression, ensuring the adjustment
234 of the model state based on observations.

235 Since each module can be executed independently, users can customise the
236 assimilation process by running only the required components, optimising
237 computational resources and allowing for seamless integration with external systems.



238 This flexibility enhances scalability, maintainability, and operational efficiency,
239 making SPREADS highly adaptable for various forecasting and research applications.

240

241 *2.3 Observation handling d4o*

242 To ensure a flexible and fast observation handling throughout all the assimilation
243 processes a query language observation database (Database for Observations, d4o)
244 based on SQLite (<https://sqlite.org/>) open-source has been developed. D4o manages
245 and controls the observations flow through observations definition and query
246 language organised in a hierarchical tree-like structure from which is easy to select
247 the desired information and place it in a data matrix for further examination. The
248 system allows SQL queries to efficiently extract and manipulate observational data.
249 The provided SQL can for example filter observations based on latitude, observation
250 type, quality control flags, and availability of posterior values. A Fortran module
251 (fd4o_mod) was developed to interface with the database. It includes functions to
252 open, close, query, and update the database. The Fortran interface is built on a C-layer
253 that calls standard SQLite APIs. The d4o database was integrated into the SPREADS
254 data assimilation system to handle large volumes of observational data. This
255 implementation optimises I/O operations and improves data accessibility across
256 parallel processing tasks. In fact, this relational-like process is particularly efficient for
257 MPI-parallel data access and queries coordination for data shuffling between MPI-
258 tasks. High vectorisation efficiency for storing and retrieving observations is therefore
259 achieved, enabling fast, flexible and configurable I/O management.

260 The Fortran interface was optimised for MPI-parallelised operations: the
261 observations are now stored and retrieved in a highly efficient, vectorised manner to
262 reduce computational overhead. Various parallelisation techniques, such as OpenMP
263 and MPI non-blocking communications and several debugging and logging options
264 were introduced to track database transactions. The database system was extended
265 to support observational inputs for different Earth system models, including CAM
266 (Community Atmosphere Model) and CLM (Community Land Model).

267 The d4o database has been architected to handle, in a fully scalable manner, the
268 vast volumes of satellite observations required for state-of-the-art analyses. It draws
269 its observation metadata directly from a series of database files (“data pools”), whose
270 contents are cached into the observational I/O-server tasks. The design of d4o



271 deliberately builds on the success of the ECMWF ODB system (Observational
272 DataBase) first deployed in 2000 to seamlessly assimilate diverse observation types
273 and, in particular, to manage very large volumes of IASI radiance data within the
274 ECMWF IFS 4DVar framework
275 ([https://www.ecmwf.int/sites/default/files/elibrary/2004/76278-ifs-](https://www.ecmwf.int/sites/default/files/elibrary/2004/76278-ifs-documentation-cy36r1-part-i-observation-processing_1.pdf)
276 [documentation-cy36r1-part-i-observation-processing_1.pdf](https://www.ecmwf.int/sites/default/files/elibrary/2004/76278-ifs-documentation-cy36r1-part-i-observation-processing_1.pdf)). In contrast to ODB, d4o
277 leverages a modern, standardised SQL-query interface via the lightweight SQLite
278 engine (<https://www.sqlite.org/>) wrapped in a versatile Fortran 2008 SPREADS
279 interface that employs hybrid MPI/OpenMP parallelism and parallel I/O for the SQLite
280 files. This hybrid approach reduces total memory usage and cuts the need for large
281 MPI task counts, avoiding the overhead of fine-grained message passing and
282 synchronisation: so that, in practice, 2–8 threads per MPI task suffice. Indeed, on Sami
283 Saarinen’s personal tests, SPREADS with its optimised d4o library running on just
284 eight nodes outperformed a comparable DART run on 32 nodes by a factor of 2–3
285 (2024–2025, pers. comm.).

286 Module-1 provides all the observations preprocessing: once all the observations
287 to be assimilated are in d4o, the screening according to the chosen resolution is
288 performed and the observations are thinned accordingly. The screening module was
289 implemented to filter out low-quality or irrelevant observations.

290 Moreover, in SPREADS, a dedicated blacklisting module has been introduced that
291 permanently excludes observation channels, platforms, or stations with documented
292 systematic errors at ingest, preventing the assimilation of problematic data and
293 reducing subsequent quality-control and computational overhead.

294

295 *2.4 Code optimisation*

296 The modifications introduced focused on optimising code efficiency and
297 enhancing parallelisation. The traditional linked-list observation sequence was
298 replaced with a SQLite-based d4o database system, significantly improving memory
299 management and data retrieval speed. An I/O server architecture was implemented
300 to separate data handling from computational tasks, reducing bottlenecks and
301 improving scalability. The sequential observation loop in Module-3 was optimised by
302 eliminating unnecessary index copying and improving observation-state closeness
303 calculations, leading to faster processing. Parallelisation was enhanced through the



304 introduction of OpenMP within MPI tasks, reducing the number of required MPI tasks
305 while maximising computational efficiency.

306 The previous blocking MPI communications were replaced with non-blocking
307 alternatives to minimise delays and improve data exchange. Additionally, database
308 operations were optimised by disabling SQLite journaling, allowing for faster
309 database writes and exclusive access for I/O servers. The handling of satellite
310 observations, particularly IASI data, was refined with a more efficient preprocessing
311 pipeline. Performance monitoring tools, such as perfstat, were introduced to identify
312 and address bottlenecks, while automated scripts were developed to streamline
313 database management, observation blacklisting, and debugging. These modifications
314 collectively enhanced the scalability, performance, and reliability of the SPREADS data
315 assimilation system. Table 1 shows the computational speed and efficiency before and
316 after the code optimisation.

Module	CPU		Node		Member	
	old	new	old	new	old	new
Module0						
Module1	105'	38'	2	1	1	1
Module2	8'	6'	12 x ts	4 x ts		
Module3	1 ^h 30	35'	25	8		

Table 1: Computer configuration and CPU time before (left panel) and after (right panel) SPREADS optimisation. Ts stands for time-slot.

317

318 **3. E-suite and preliminary results**

319 SPREADS is fully integrated into the Community Earth System Model (CESM,
320 <https://www.cesm.ucar.edu/>), an infrastructure developed through a joint
321 collaboration between many meteorological centres including CMCC

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328 and NCAR. CESM provides a flexible software framework for configuring and

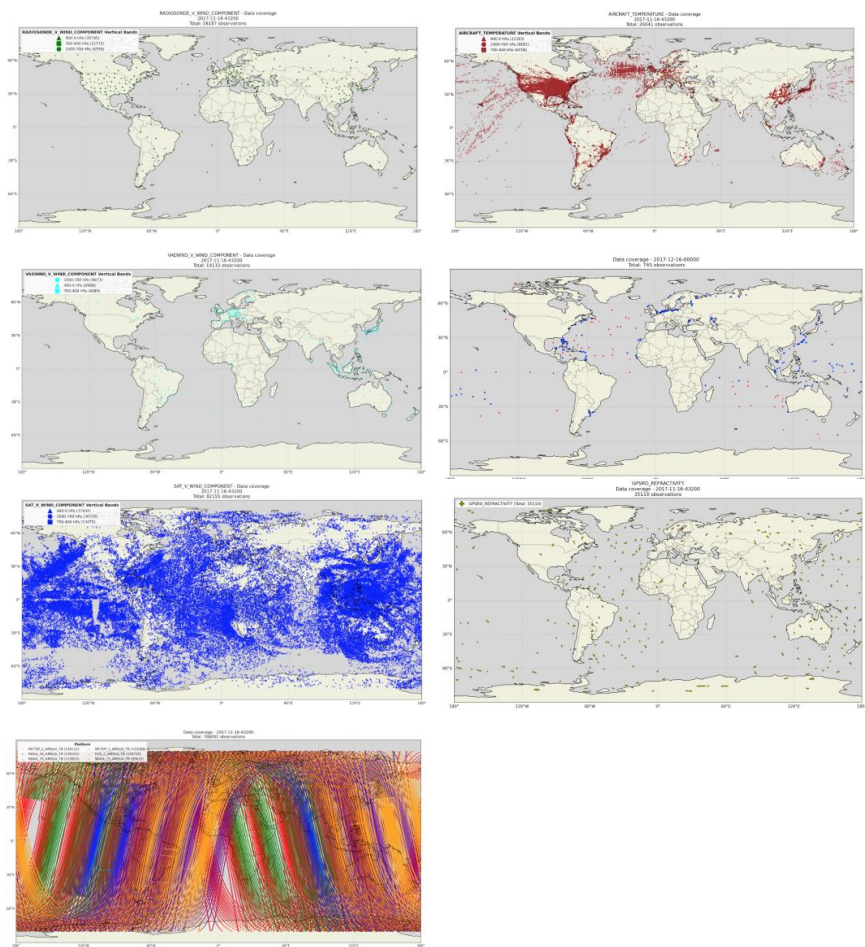


Figure 2: 6 hour window observation data coverage according to observation type. From top left: Radiosonde, Aircraft, Wind profiler, Synop and Buoy, AMV, GPS-RO and AMSU-A

329 running coupled models, each designed to represent different components of the
330 Earth system. Specifically, SPREADS is coupled with the atmosphere model (CAM), the
331 land model (CLM), the cryosphere model (CICE), and the ocean model NEMO (for the
332 CMCC CESM).
333



Observation	Observation kind	Information
GPS-RO Metop A, B GRAS	Refractivity	Temperature
AMSU-A Metop A	Microwave sounder radiance	Temperature
AMSU-A Metop B		
AMSU-A NOAA-15		
AMSU-A NOAA 18		
AMSU-A NOAA-19		
AMSU-A AQUA		
AMV AQUA	Visible	u, v
AMV TERRA	Visible	
AMV GOES-15	Visible	
AMV Meteosat-10	IR, WV and V	
AMV COMS-1	IR, WV and V	
AMV Dual Metop	IR	
AMV INSAT -3D		
AMV NPP	IR	
AMV HIMAWARI-8	Visible	
Profiler	European, Japanese Wind	u, v
Radiosonde	Land and Ship	u, v, T, q
Aircraft		u, v, T, q
Buoys	Moored and Drifters	Surface pressure
SYNOP	Land	Surface pressure

Table 2 Observation types and platforms assimilated

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335 An experimental suite (E-suite) has been implemented using SPREADS to assess
336 its performance in an operational-like environment. The E-suite began running in
337 January 2024, covering the period from July 2017 to the present at 1° horizontal
338 resolution. It is currently up to the January 2018 analysis production, utilising the 93
339 model levels of CAM (version 6 finite volume dynamical core; Simpson *et al*, 2025),
340 with enhanced vertical resolution in the free troposphere and stratosphere and a



341 model lid height set at 0.01 hPa. Each cycle operates on a 6-hour window, subdivided
 342 via the FGAT approach to accommodate the asynchronous nature of the observations.
 343 The E-suite assimilates a wide range of observations, as listed in Table 2.
 344 A representative data coverage for a 6-hour assimilation window centred on 12
 345 UTC is shown in Figure 2. For that cycle, the number of assimilated observations

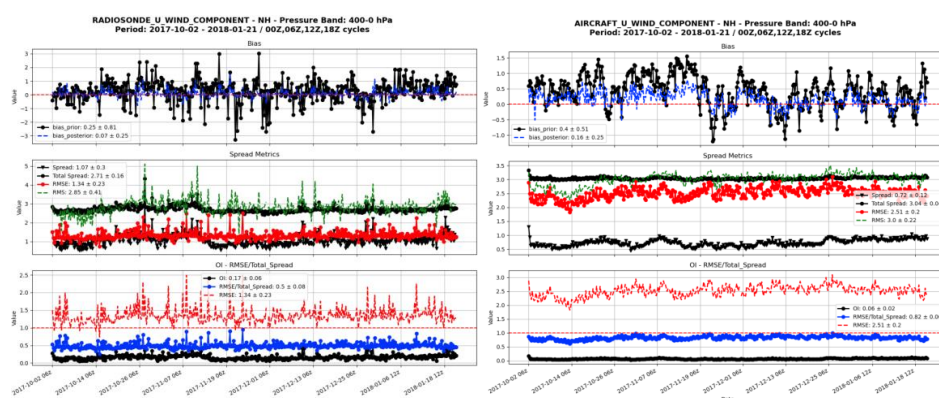


Figure 3: Time series of observation-minus-background (prior) and observation-minus-analysis (posterior) departures for the zonal wind component (u) from 2 October 2017 to 21 January 2018 at 00, 06, 12, and 18 UTC, over the Northern Hemisphere and between 400 hPa and 0.01 hPa. The left panel shows radiosonde observations, and the right panel shows aircraft observations. Displayed diagnostics include RMS, RMSE, ensemble Dispersion Index (DI), and Observation Influence (OI)

346 includes: 35,110 GPS-RO refractivity profiles, 768,002 AMSU-A brightness
 347 temperatures, 163,284 AMV winds, 40,158 wind profiler reports, 174,263 radiosonde
 348 measurements, 79,043 aircraft reports, and 500 surface pressure observations from
 349 SYNOP and BUOY platforms. This results in a total of approximately 1.3×10^6
 350 assimilated observations every 6 hours. Figure 3 presents time series of the zonal
 351 wind component (u) from radiosonde (left) and aircraft (right) observations over the
 352 Northern Hemisphere between 400 hPa and 0.01 hPa. The assimilation results show
 353 a reduction in bias of approximately 30% for radiosondes and 40% for aircraft data.
 354 Aircraft observations typically exhibit smaller biases, generally within ± 1.5 m/s,
 355 compared to radiosonde data, which show biases reaching ± 3 m/s.



356 A notable change occurs around 14 November 2017, coinciding with the
 357 introduction of AMSU-A microwave radiances into the assimilation system; prior to
 358 this, aircraft data tend to underestimate the zonal wind.

359 The RMSE for aircraft observations remains higher (~ 2.5 m/s) than that for
 360 radiosondes (~ 1.5 m/s), consistent with known instrument characteristics and
 361 sampling differences. The ensemble Dispersion Index (DI), defined as the ratio of the
 362 ensemble RMSE (computed against the own analysis) to the total spread, stays close
 363 to 1 throughout the period for the aircraft observation types, whilst is less than 1 for
 364 the radiosonde observations indicating an overdispersive ensemble at the top of the
 365 atmosphere (~ 50 hPa) and a well calibrated ensemble in the high troposphere (~ 250
 366 hPa).

367 The observation Influence (OI), $0 \leq OI \leq 1$, indicates that when $OI = 0$ the
 368 observation had no leverage in the fit, whereas $OI = 1$ means the fit relied entirely on
 369 the observation, with no contribution from the first guess (Cardinali *et al.*, 2004;
 370 Cardinali, 2014; Liu *et al.*, 2009; Gharamti *et al.*, 2019). OI remains relatively low for
 371 aircraft (~ 0.1), while radiosondes display higher influence values (~ 0.25), reflecting
 372 their higher information content and smaller observation error variances in the upper
 373 troposphere and lower stratosphere.

374 AMSU-A channels 9–14 were assimilated starting on 14 November 2017, initially
 375 with a scan bias correction following Harris and Kelly (2001) and subsequently, on 2
 376 January 2018, with the addition of an air-mass bias correction following Noh *et al.*
 377 (2023). While the scan bias correction, after extensive evaluation, was shown to
 378 perform satisfactorily, the air-mass bias correction proved less effective.

379 Further analysis demonstrated that the regression predictors used for the 200–
 380 50 hPa thickness were inadequate to represent channels 11–14. A more suitable
 381 choice was to separate the predictors and also include a 50–2 hPa thickness to account
 382 for the stratospheric channels. Therefore, the final air-mass bias correction was
 383 applied using a linear combination of several thickness predictors (1000–300, 200–
 384 50, and 50–2 hPa), enabling the scheme to account for biases arising from multiple
 385 physical dependencies simultaneously (Auligné *et al.*, 2007). To anchor the model

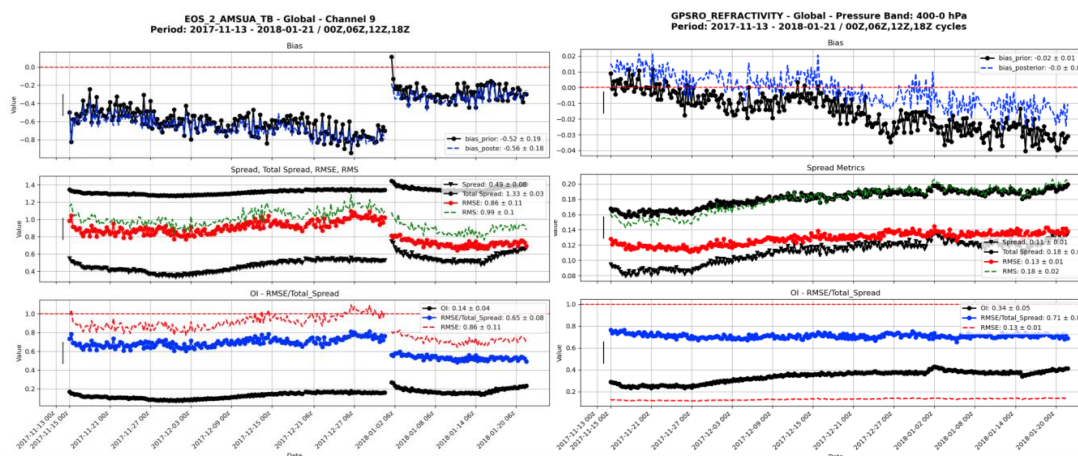


Figure 4: Global time series of Aqua AMSU-A TB channel 9 (left panel) compared with GPS-RO refractivity in the layer 400-0.01 hPa (right panel).

386 bias, unbiased radiosonde and GPS-RO observations were used, and AMSU-A channel
 387 14 was left uncorrected since it is considered unbiased for the same reason. In
 388 addition, the averaging of correction coefficients across last four cycles, as done in Noh
 389 *et al.*, was found to mask the actual dynamical situation. To better capture flow-
 390 dependent variability, our implementation used only the predictors within the current
 391 assimilation window for the regression.

392 Figure 4 presents the time series diagnostics for two key observation types, GPS-
 393 RO refractivity (right) and AMSU-A AQUA Channel 9 brightness temperatures (left),
 394 covering the global pressure range from 400 to 0 hPa over the period 20171113 to
 395 20180122. The GPS-RO diagnostics indicate a well-calibrated assimilation system: the
 396 posterior bias remains near zero (0.00 ± 0.01), the RMS and total spread are closely
 397 matched, and the dispersion index remains near one. The observation influence (OI)
 398 is moderate (0.34 ± 0.05), showing a balanced contribution between the observations
 399 and the model background. These results confirm GPS-RO's role as a high-impact,
 400 high-precision observation source, especially in the upper troposphere and
 401 stratosphere.

402 In contrast, AMSU-A Channel 9, which peaks near 100 hPa in the mid-
 403 stratosphere, presents a more complex picture. Following the introduction of scan
 404 angle bias correction on 14 November and air-mass bias correction on 2 January, a



405 notable reduction in posterior bias is observed from ~ 0.6 K to -0.2 K. This
 406 improvement in bias is accompanied by a steady decrease in RMSE from ~ 1 to 0.8 K.
 407 Additionally, the observation influence rises to a higher value of 0.25 , indicating a
 408 better leveraged analysis fit.
 409

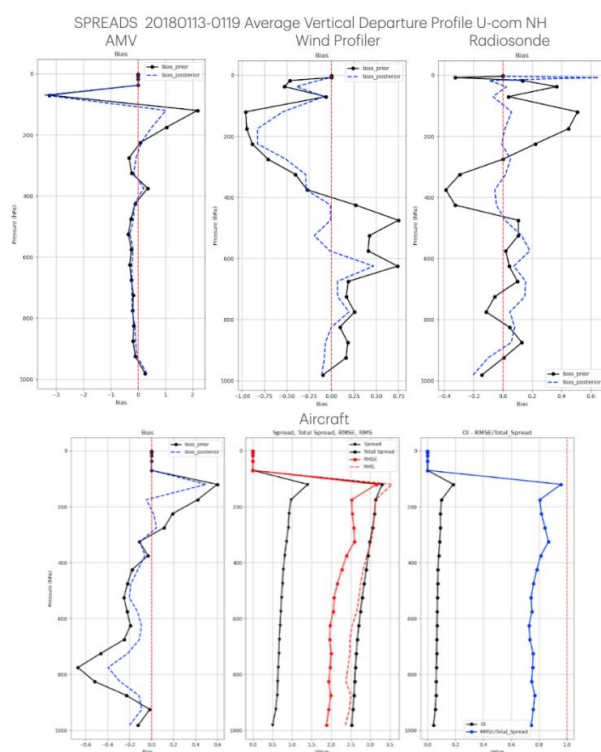


Figure 5: Average vertical profiles of the zonal wind component (u) for AMVs, wind profilers, and radiosondes (top panel), and for aircraft observations (bottom panel). Aircraft profiles are complemented by diagnostics including RMS, RMSE, Dispersion Index (DI), and Observation Influence (OI).

410 This diagnostic reinforces conclusions that while SPREADS effectively assimilates
 411 high-accuracy GPS-RO data, stratospheric radiance assimilation still presents some
 412 challenges.

413 Figure 5 presents a comparison of the prior and posterior departure vertical
 414 profiles of the zonal wind, averaged over the week 20180103–19, for AMV, Wind
 415 Profiler, and Radiosonde observations (top panel), and Aircraft observations (bottom



416 panel). For the Aircraft data, ensemble performance statistics are included: ensemble
 417 spread, total spread, RMS, RMSE, dispersion index, and observation influence (OI).

418 Across all observation types, the posterior fit demonstrates a clear reduction in
 419 bias throughout the atmosphere. Larger residual departures are seen above 200 hPa,
 420 with AMVs showing differences between -3 m/s at 50 hPa and $+1$ m/s at 100 hPa.
 421 Wind Profiler data shows a -0.75 m/s departure at 200 hPa, while Aircraft exhibit
 422 slightly positive departures of 0.4 m/s at around 100 hPa. These discrepancies likely
 423 stem from the sparser data coverage at upper levels and some residual effects from
 424 AMSU-A assimilation.

425 The profiles also show that posterior fits tend to converge across observing
 426 systems in the troposphere, suggesting a consistent adjustment by the assimilation
 427 system despite differing data characteristics. A subtle transition in departure
 428 behaviour is visible near the tropopause, possibly indicating increased
 429 representativeness error or limitations in vertical resolution at this level.

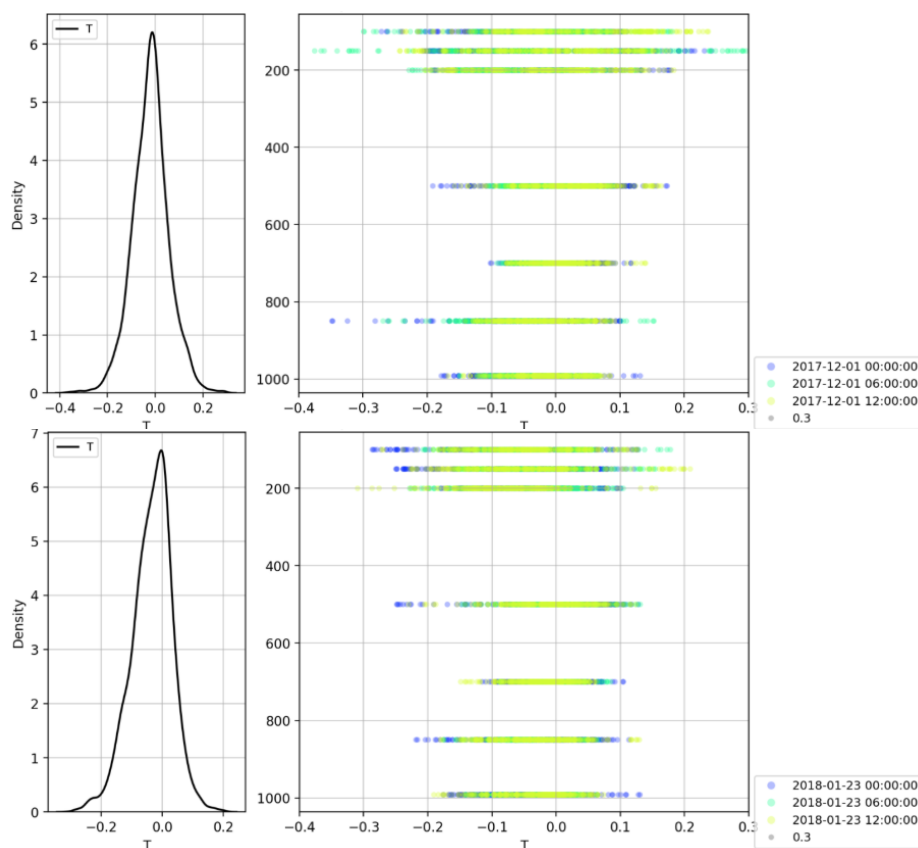


Figure 6: Temperature increment density distribution for 20171201 (top panel) and 20180123 (bottom panel). Probability density function of the increments (left) and their vertical distribution across 3 cycles (right)



430 Ensemble performance metrics are shown only for Aircraft data, as their
 431 behaviour is representative of the other platforms. The RMSE remains around 2 m/s
 432 throughout the column, while the RMS departures (middle panel, bottom row) are
 433 slightly higher, as expected. The dispersion index indicates approximately 20% over-
 434 dispersion in the ensemble below 200 hPa. The OI for Aircraft observations is
 435 relatively low (0.1 to 0.2), primarily due to their error characteristics.

436 Finally, since the statistics are averaged over a full week, they represent
 437 systematic patterns rather than short-term variability. Interpretation of vertical
 438 features should also consider the varying vertical resolution and density of each
 439 observation type, particularly above 200 hPa, where reduced data availability can
 440 influence both bias correction and ensemble reliability.

441 Figure 6 illustrates the temperature increment density distribution for two
 442 selected dates during the analysis period: December 1st, 2017 (top panel) and January
 443 23rd, 2018 (bottom panel). Each panel shows the probability density function of the
 444 increments (left) and their vertical distribution across three analysis cycles of each
 445 day (right).

446 Over the course of the period, a clear reduction in the amplitude of temperature
 447 increments is observed, particularly within the troposphere. The spread of the
 448 distribution narrows from approximately $\pm 0.4^\circ\text{K}$ on December 1st to about $\pm 0.2^\circ\text{K}$ by
 449 January 23rd. This contraction reflects improved constraint in the analysis, likely
 450 resulting from better-calibrated observations and/or enhanced ensemble
 451 performance.

452 Alongside the amplitude reduction, the tropospheric increment distributions
 453 become increasingly symmetric around zero, indicating a progressive reduction in
 454 systematic bias and a more balanced assimilation system. Additionally, the analysis
 455 cycles per day show greater consistency over time: while 20171201 displays notable
 456 variability between cycles, especially in the lower troposphere, the cycles on
 457 20180123 exhibit much tighter agreement, suggesting improved temporal stability of
 458 the system. Similar reduction of the increments is observed in the stratosphere.

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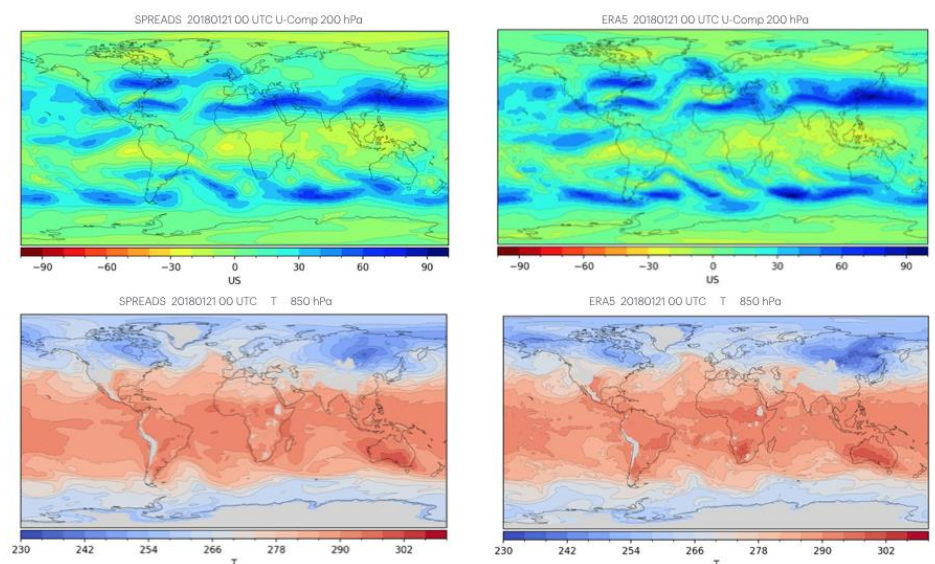


Figure 7: SPREADS analysis (left panels) compare with ERA5 (right panels) of the U wind component at 200 hPa (top panels) and T at 850 hPa (bottom panels)

462 Figures 7 and 8 present a comparison between SPREADS and ERA5 analyses,
 463 highlighting both horizontal and vertical structural differences in the representation
 464 of key atmospheric variables.

465 Figure 7 displays global fields from SPREADS (left panels) and ERA5 (right
 466 panels) at 00 UTC on 20180121 for the zonal wind component at 200 hPa (top) and
 467 temperature at 850 hPa (bottom). The large-scale circulation patterns are well
 468 captured in SPREADS, showing good agreement with ERA5 in both magnitude and
 469 spatial structure. However, differences emerge at smaller spatial scales, particularly
 470 in regions with sharp gradients such as subtropical jet streams, where ERA5 exhibits
 471 finer, more coherent jet streaks owing to its higher horizontal resolution nearly
 472 double that of SPREADS.

473 At 850 hPa, the temperature patterns in both systems reflect realistic meridional
 474 gradients and contrast between land and ocean. Yet, subtle regional differences are
 475 visible: SPREADS appears cooler over high-latitude

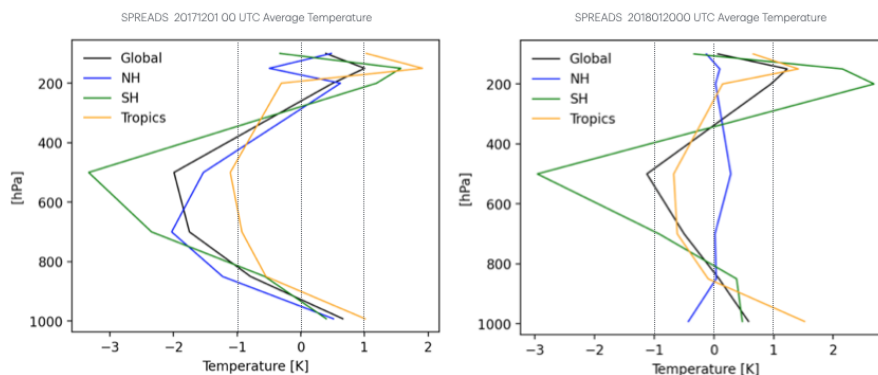


Figure 8: Vertical profile of the average T differences between SPREADS and ERA5 for the Global (black line), NH (blue line), SH (green line) and TR (yellow line) valid at 00 UTC 20171201 (left panel) and 20180121 (right panel)

476 continental regions, such as Siberia and Canada, possibly due to differences in
 477 land-surface model physics or less dense observational constraints in those areas. The
 478 thermal contrast between land and ocean is well maintained in both datasets, though
 479 slightly smoother in SPREADS, again reflecting resolution effects.

480 To further understand these differences, Figure 8 shows vertical profiles of
 481 average temperature differences (SPREADS minus ERA5) at 00 UTC for two key dates:
 482 20171201 (left) and 20180120 (right), representing the beginning and end of the
 483 evaluation period. The profiles are shown for Global (black), Northern Hemisphere
 484 (blue), Southern Hemisphere (green), and Tropics (orange).

485 From early December to late January, the NH and Tropics show a marked
 486 improvement, with temperature differences decreasing by up to ± 1 K, particularly in
 487 the mid-to-upper troposphere. This reflects both improved background constraint
 488 and effective assimilation updates during the SPREADS evaluation period. In contrast,
 489 the SH shows minimal change, likely due to sparser observational coverage,
 490 underscoring the asymmetry in observing system density between hemispheres.

491 Notably, around 200–300 hPa, the NH and Tropics exhibit a transition from cold
 492 to warm bias, suggesting a tropopause-level sensitivity that may be

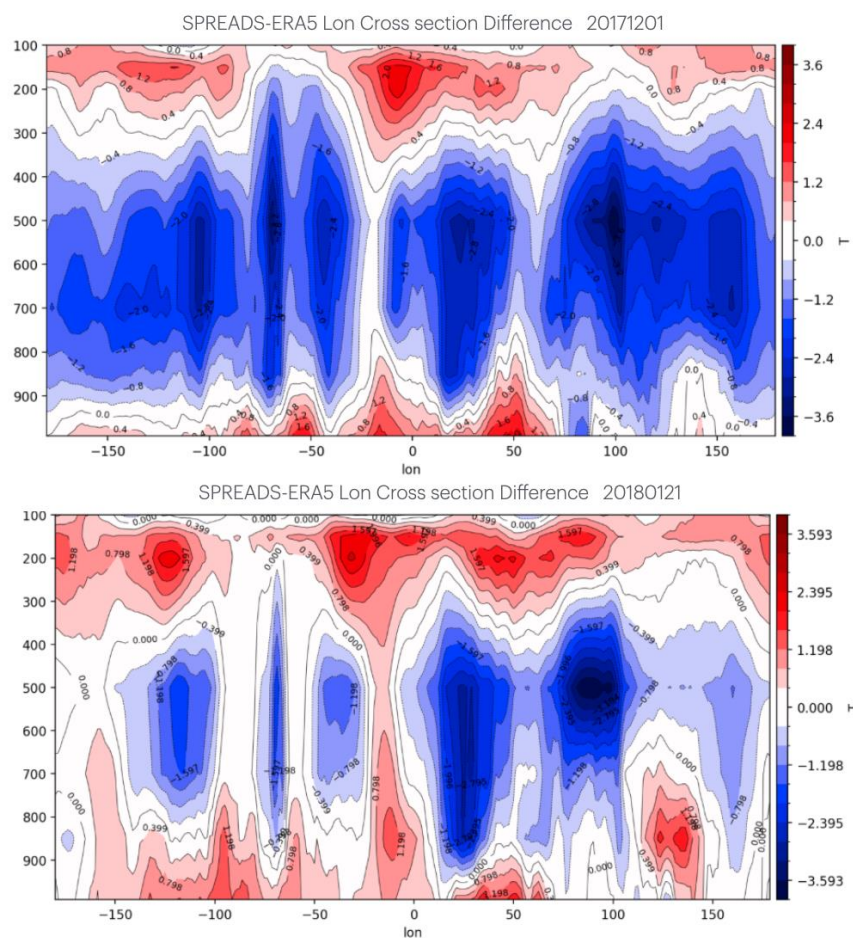


Figure 9: Zonal cross section of the T differences between SPREADS and ERA5 valid at 00 UTC 20171201 (top panel) and 20180121 (bottom panel)

493 influenced by the vertical resolution and radiative transfer modelling. Above 100,
 494 a persistent warm bias remains globally, more pronounced in the SH, pointing to
 495 potential limitations in stratospheric observation assimilation, possibly related to
 496 reduced usage of GPS-RO or upper-level radiance channels. Overall, despite the
 497 coarser resolution, SPREADS reproduces the dominant features of the atmospheric
 498 state with high fidelity. Improvements over time in the troposphere, particularly in
 499 the NH and Tropics, indicate that the system is maturing well. Future enhancements,
 500 such as expanded use of satellite data in the SH, and increased spatial resolution, could
 501 further close the gap with ERA5 in under-constrained regions.



502 As a final examination of the analysis difference, Fig. 9 and 10 show the average
 503 cross section of temperature for the E-suite initial day (top panel) and at the end of
 504 January (bottom panel) across the longitude and latitude, respectively. At the initial
 505 time (Fig. 9 top panel), the temperature differences between SPREADS and ERA5 show
 506 widespread cold biases (blue shading) across nearly all longitudes in the lower to mid-
 507 troposphere (around 900–400 hPa). Maxima of positive differences exceed +3.6 K,
 508 particularly around the central and eastern Pacific and parts of the Atlantic. In the
 509 upper troposphere (above ~300 hPa), some warm biases (red shading) begin to
 510 appear, though they are less dominant. Figure 9 bottom panel shows a clear
 511 improvement in the zonal consistency and magnitude of the differences. Cold biases
 512 are notably reduced in amplitude and spatial extent, particularly across the mid-
 513 troposphere. The structure becomes more vertically layered and less zonally
 514 coherent, suggesting improved local balance and constraint. Warm anomalies aloft
 515 become slightly more pronounced in some sectors (e.g., near 60° longitude), pointing
 516 to evolving differences in vertical structure, possibly due to the upper-air
 517 observational influence. The complemented Fig. 10 initially shows (top panel)
 518 significant cold biases (~-3 to -6 K) in the extratropics, especially over the Southern
 519 Ocean (around 60°S–80°S) and Northern Hemisphere high latitudes (~60°N–80°N),
 520 spanning from the surface up to 400 hPa. The tropical region remains relatively
 521 neutral, with near-zero or weakly positive anomalies. The biases show a strong
 522 hemispheric asymmetry, being more intense and vertically extensive in the SH. At the
 523 final time (bottom panel), the magnitude of cold biases is substantially reduced,
 524 particularly over the Southern Hemisphere, where mid- and upper-tropospheric
 525 differences nearly vanish. Remaining differences are more localised and patchy, with
 526 some persistent warm
 527 anomalies in the upper troposphere (above 300 hPa) over both poles. The tropics
 528 remain stable, with minor variations and low-magnitude anomalies. There is an
 529 overall flattening of the difference structure, indicating improved vertical and
 530 hemispheric balance. In summary, between early December and late January,

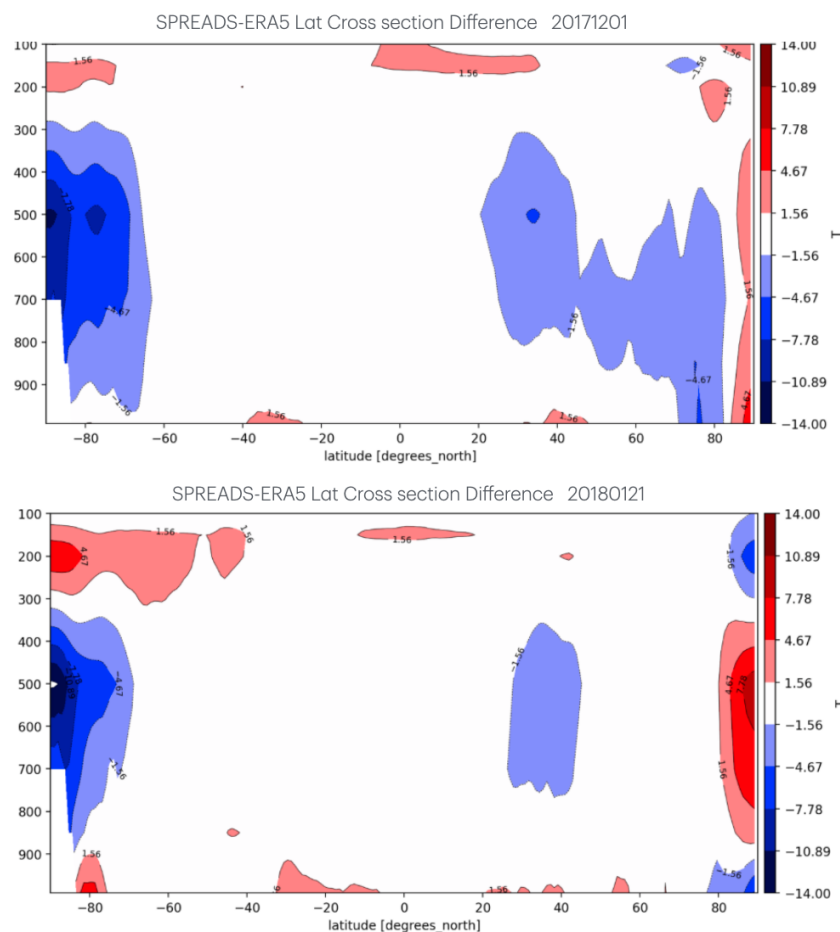


Figure 10: Meridional cross section of the T differences between SPREADS and ERA5 valid at 00 UTC 20171201 (top panel) and 20180121 (bottom panel)

SPREADS shows significant improvement in temperature alignment with ERA5, particularly in the lower to mid-troposphere and across the Southern Hemisphere. The initial large cold anomalies in the SH and NH extratropics largely diminish, suggesting better background constraint or improved assimilation tuning over time. There is a slight emergence of warm biases in the upper troposphere and lower stratosphere, possibly tied to model vertical resolution (Simpson *et al.*, 2025) or under observed stratospheric layers. In general, the differences become less globally coherent and more structured, indicating that the system is moving toward finer-scale, observation driven corrections rather than broad model biases.

4. Conclusion&Plan



543 The development of SPREADS, Scalable PaRallelised EArth Data Assimilation
544 System, represents a crucial step in advancing ensemble-based data assimilation from
545 research to operational application. Built upon the flexible, open-source DART
546 framework, SPREADS embodies the principles of transparency, collaboration, and
547 reproducibility that are foundational to modern Earth system science. Open-source
548 modelling is not only a technical choice but a strategic enabler of scientific progress:
549 it fosters community-driven innovation, ensures the traceability of results, and
550 accelerates the adoption of new ideas across institutions and research domains. By
551 sharing tools, code, diagnostics, and configuration options, SPREADS positions itself
552 at the forefront of a collaborative data assimilation system.

553 SPREADS introduces a suite of technical enhancements, FGAT-based temporal
554 alignment, modular parallelised assimilation architecture, and the d4o SQL-based
555 observational database to address the computational and algorithmic challenges of
556 operational ensemble systems. These advancements enable the efficient assimilation
557 of over one million observations every six hours, including a diverse set of
558 conventional and satellite-based measurements. The E-suite evaluation demonstrates
559 promising results, with improved bias characteristics, ensemble calibration, and
560 overall consistency in comparison to ERA5, particularly in the Northern Hemisphere
561 and tropics.

562 The bias diagnostics in SPREADS are consistent with model-based findings from
563 the CAM7 vertical resolution study (Simpson *et al.*, 2025). In particular, the cold biases
564 observed in the tropical lower stratosphere in SPREADS, as diagnosed through AMSU-
565 A, align with those seen in low-vertical-resolution CAM configurations. The
566 application of FGAT and adaptive bias correction in SPREADS significantly reduces
567 these biases, mirroring the improvements achieved in CAM7 through enhanced
568 vertical resolution. This convergence from both model and observational assimilation
569 perspectives underscores the robustness of the diagnostic framework in SPREADS
570 and its capability to detect and mitigate systematic biases in the upper troposphere
571 and lower stratosphere. A key direction for SPREADS development is the inclusion of
572 more satellite-based observations, particularly infrared radiances such as IASI. These
573 sensors provide rich vertical information in cloud-free conditions and are essential for
574 improving temperature and humidity profiles, especially in the stratosphere and
575 upper troposphere. Preliminary testing of IASI data within SPREADS is currently



576 underway and shows great promise for enhancing the vertical structure of the
 577 analysis and addressing residual biases observed in the current system.

578 In parallel, all-sky microwave radiance assimilation is being actively tested. This
 579 represents a shift in satellite data usage, enabling the assimilation of radiances under
 580 both clear and cloudy conditions. All-sky assimilation significantly increases the
 581 spatial and temporal coverage of radiance data, especially in regions with persistent
 582 cloud cover such as the tropics and storm tracks. By more effectively capturing cloud-
 583 affected observations, SPREADS aims to improve its representation of moisture fields,
 584 cloud dynamics, and convective processes, key elements for accurate medium- to
 585 long-range forecasts.

586 Although SPREADS currently operates at coarser resolution than ERA5, its ability
 587 to replicate large-scale atmospheric patterns and to reduce biases over time
 588 demonstrates the system's robustness. Continued tuning of satellite bias corrections,
 589 expansion of satellite data types, and enhancement of vertical resolution will be
 590 critical next steps. Furthermore, ongoing integration within a fully coupled Earth
 591 system model positions SPREADS as a strategic asset for seamless forecasting, from
 592 weather to climate timescales.

593 In conclusion, SPREADS is a scalable, open, and forward-looking platform that
 594 effectively bridges research innovation with operational demands. Its modular,
 595 transparent architecture invites community contribution and ensures adaptability to
 596 evolving scientific goals. With a growing observational portfolio and expanding
 597 capabilities, including the assimilation of all-sky and hyperspectral infrared radiances,
 598 SPREADS is well positioned to become a next-generation system for global Earth
 599 system prediction.

600

601 **Code and data availability**

602 All codes in this study are permanently available at
 603 <https://doi.org/10.5281/zenodo.17063454> (Cardinali *et al.*, 2025).

604

605 **Author contribution**

606 CC: Conceptualisation, Methodology, Supervision, Project administration,
 607 Validation, Investigation, Writing original draft.

608 GC: Methodology, Code developments, Investigation



609 MG: Infrastructure development
610 SS: Observation database development and Code optimization
611 GD: Code developments
612 JA: Methodology and Code development support
613 KR: Software and Technical support
614

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616

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