



Rapid Landslide Mapping During the 2023 Emilia-Romagna Disaster: Assessing Automated Approaches with Limited Training Data

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Abstract

The catastrophic rainfall events of May 2023 in the Emilia-Romagna region, Italy triggered more than 80.000 landslides (<https://doi.org/10.5281/zenodo.13742643>, Pizziolo et al., 2024) and placed extraordinary demands on emergency response systems. One of the most critical tasks during the emergency was mapping the landslides caused by the event, a process carried out manually that required substantial time and effort (Berti et al., 2024). This study explores the potential of automated landslide mapping to support rapid disaster response, evaluating two deep learning models, U-Net and SegFormer, under realistic emergency constraints, including minimal training data. The models were trained exclusively in one affected municipality area (Casola Valsenio, 84 km²) and tested on three additional municipalities (Predappio, 91 km²; Modigliana, 101 km²; Brisighella, 194 km²) with varying geological settings. To reflect a range of operational scenarios, we tested seven combinations of input layers, progressively increasing in complexity from post-event Sentinel-2 imagery alone to full integration of high-resolution aerial imagery, NDVI change maps, and slope data.

Results show that both models achieved comparable segmentation performance, with SegFormer displaying greater robustness to variations in input layers and geological conditions, while U-Net was more sensitive but occasionally more accurate with rich inputs. Both models successfully identified landslides, but encountered difficulties in shadowed zones, cultivated fields, and geologically distinct terrains. A key limitation emerged in Brisighella FAA, a sector dominated by Blue Clay formations, where poor generalization was traced to the lack of lithological diversity in the training data, highlighting the need for geologically balanced training datasets.

Despite these challenges, the study confirms the operational value of automated mapping as a first-pass tool. Both models delivered accurate baseline maps that could support manual validation and prioritization in time-critical scenarios. The findings suggest that AI-based mapping could become a key component of emergency management protocols. While manual revision remains essential, these tools offer a scalable, time-efficient solution to the increasing demand for rapid and spatially detailed hazard information.



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41 1. Introduction

42 Rapid and accurate landslide mapping over large areas is a challenging task, particularly in the context
43 of regional-scale rainfall or earthquake events (Iverson et al. 2015; Casagli et al. 2016; Robinson et
44 al., 2017; Holbling et al. 2017; Mondini et al., 2021). This task is crucial for implementing timely and
45 effective response measures, assessing the extent of impact, and planning for recovery and mitigation
46 efforts. The complexities of rapid landslide mapping vary depending on several factors including the
47 extent of the affected area, the availability of cloud-free imagery, the precision needed for the
48 mappings, and the difficulties of detecting landslides due to shadows, vegetation cover, or weak
49 geomorphic evidences (Mondini et al. 2019; Amatya et al., 2023). Typically, these complexities are
50 addressed by expert geologists through the visual interpretation of satellite or aerial imagery (Guzzetti
51 et al., 2012; Scaioni et al., 2014; Ferrario and Livio 2023). These experts are skilled at identifying
52 subtle variations in the terrain and can integrate complex contextual knowledge, including field
53 reports, personal accounts, and specific geological information, to create highly accurate landslide
54 maps. However, manual mapping is both time-consuming and inherently subjective, which are
55 significant limitations during a crisis (Novellino et al. 2024).

56 Automated landslide recognition techniques using convolutional neural networks (CNN) algorithms
57 are emerging as promising alternatives to manual methods (Sameen and Pradhan 2019, Chen et al.
58 2018, Ye et al. 2019, Tang et al. 2022, Ji 2020). Recent studies have demonstrated their effectiveness
59 in real-world cases. For instance, Meena et al. (2021) applied a deep learning approach for rapid
60 landslide mapping in India following extreme monsoon rainfall, showcasing the potential of these
61 techniques in post-disaster conditions. Prakash et al. (2021) introduced a new strategy using a
62 generalized convolutional neural network, aiming to develop a more versatile method applicable
63 across various geographical contexts. Building on this, Prakash and Manconi (2021) demonstrated
64 the practical utility of these methods by rapidly mapping landslides triggered by a severe storm event.

65 These studies emphasize the growing importance of CNNs in automated landslide mapping and their
66 potential for practical implementation in disaster management. However, despite their proven
67 effectiveness, a notable gap remains between theoretical applications and their real-world deployment
68 during emergencies. Typically, these studies employ a large train-to-validation ratio (usually 80:20
69 or 70:30; Meena et al 2023), which does not mirror the limited data availability for training models
70 in real-time crisis situations. Furthermore, the training process frequently does not account for the
71 variety of landslide types and the spectrum of geological conditions present in actual situations. This
72 oversight is critical, as the complexity of natural environments play a crucial role in the effective
73 application of these technologies in emergency contexts.

74 In this study, we utilize the landslide inventory from the May 2023 crisis in the Emilia-Romagna
75 region of Italy (Berti et al. 2025) to assess the effectiveness and limitations of automated mapping
76 algorithms in a practical setting. We specifically focus on the performance of the U-Net and
77 SegFormer algorithms in identifying landslides triggered by the event, despite the challenges of very
78 limited training data. The study also considers the variety of landslide types and geological conditions
79 prevalent in the area, which impacted the disaster response. The main goal of this research is to
80 determine whether advanced deep-learning methods can effectively replace manual mapping in



managing large-scale landslide disasters, thus bridging the gap between academic research and real-world application in emergency management.

2. The Romagna May 2023 disaster

2.1 Overview of the disaster

In May 2023, the Emilia-Romagna region of Italy was hit by unprecedented meteorological events, as detailed by Berti et al. (2025) and Foraci et al. (2023). Two major rainfall episodes, from May 1-3 and May 16-17, led to severe floods and landslides primarily affecting the eastern part of the region (Fig. 1a). The first event (May 1-3) was driven by a stationary low-pressure system which delivered over 200 mm in 48 hours. The second event (May 16-17), featured a similar low-pressure system enhanced by bora winds and moist air from the Mediterranean, resulting in rainfall exceeding 250 mm in the same duration. The cumulative impact of these episodes resulted in more than 500 mm of rainfall over two weeks, nearly half the region's annual climatic precipitation (Fig. 1a). The return period for these combined rainfall events was estimated to exceed 1000 years (Brath et al. 2023), highlighting the unprecedented nature of this occurrence. The two cyclonic events led to extensive flooding, submerging 540 km² of land, and triggered more than 80,000 landslides across approximately 1000 km² (Fig. 1b). This widespread devastation impacted thousands of buildings, roads, and other infrastructure, with the estimated damages amounting to approximately 9 billion euros.

Immediately after the disaster, we initiated the development of a detailed inventory map of the landslides caused by the event. This map was manually created using high-resolution aerial imagery taken only one week following the second rainfall event. The images, which include four bands (RGB+NIR) and have a resolution of 0.2 m, allowed for the accurate identification and classification of the landslides. The data derived from this mapping effort are described in Berti et al. (2025) and are available for public download in shapefile format (DOI: 10.5281/zenodo.13742643, Pizziolo et al., 2024).

2.2 Motivation of the work

The visual identification of the landslides from May 2023 proved to be a challenging and time-intensive process. It took six months for twelve expert geologists to identify and delineate all the landslides within the 1000 km² area affected by the disaster. Despite these challenges, this approach was deemed the best available method at the time to meet the Civil Protection agency's specific requirements for accuracy, precision, and oversight during the emergency response coordination.

As this manual mapping progressed, we explored various automated mapping techniques to speed up the process and reduce the workload. Initial trials using NDVI methods and U-Net in a specific sample area produced promising results, demonstrating that deep-learning models can effectively replicate manual mapping efforts (Berti et al., 2024). However, these preliminary analyses were conducted in a test area (the Casola Valsenio municipality, Fig. 1), which covers only about 10% of the total area impacted by the event. Consequently, several critical questions remain unanswered: i) Can automated models trained on a small area be effectively used for rapid mapping of larger regions? ii) How do



models trained in specific geological conditions perform across different geological settings? iii)
Could more sophisticated deep-learning methods surpass U-Net in improving outcomes?
Addressing these questions is crucial to determine whether emerging automated methods can enhance
the efficiency of landslide mapping in future crisis scenarios, thereby potentially reducing critical
response times. In this study, we simulate landslide mapping during a crisis, taking into account the
limited data typically available and the accessible information under such conditions. This simulation
is grounded in our firsthand experience of the real emergency during the May 2023 event, which
provided us with valuable insights into the practical challenges and data limitations encountered
during such crises.

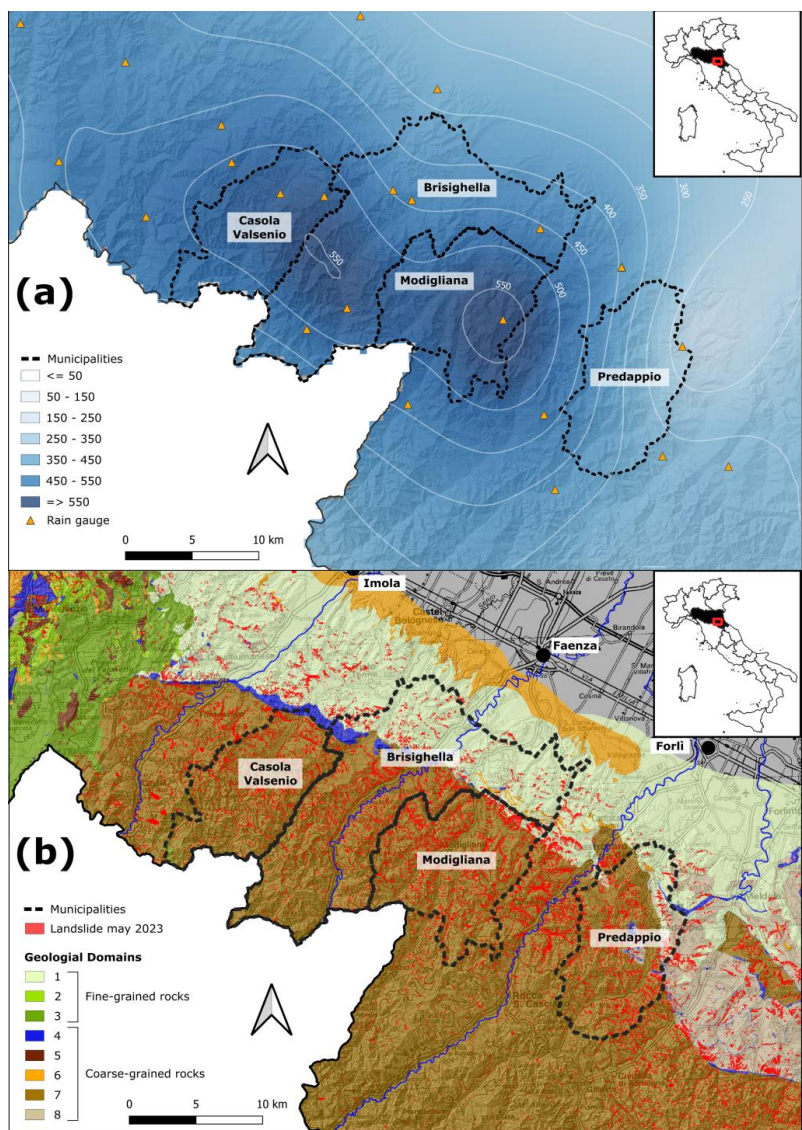


Figure 1. Overview of the Effects of the Extreme Rainfall Event in Romagna, May 2023.



a) Cumulative precipitation map from May 1 to May 17, 2023. b) Manually mapped landslide distribution across the study areas, categorized by geological formations as per the classification by Berti et al. (2025).

3. Study area

Automated methods for landslide mapping were applied in the four municipalities shown in Figure 2. These areas, situated in the eastern part of the Emilia-Romagna region (Fig. 1), were significantly affected by the meteorological event of May 2023. The Casola Valsenio area was chosen for training models, whereas Brisighella, Modigliana, and Predappio served as validation sites. In their study, Berti et al. (2024) used Casola Valsenio to train a conventional U-Net model, and compared the results with those obtained using the traditional NDVI change detection method. We opted to continue using this area for training to ensure consistency with previous research. The other three areas were selected because together they encompass approximately 35% of the regions most impacted by landslides and represent about 25% of the recorded 80,997 landslide events. Moreover, these municipalities exhibit varied geological conditions, with both similarities and differences to Casola Valsenio, thereby offering a comprehensive test of the models' adaptability to diverse terrains and landforms.

To represent geological variability in a way that is both meaningful and operational for the analysis, we adopted the lithological classification introduced by Berti et al. (2025), which groups the over 600 geological formations of the Emilia-Romagna region into eight lithological units (Fig. 1b). This classification combines lithological characteristics with structural domains, acknowledging that the same rock type may exhibit different mechanical properties depending on its tectonic setting. Units 1 to 3 consist of fine-grained materials (e.g., clays, marls, tectonized shales), while units 4 to 8 mainly comprise coarse-grained rocks (e.g., sandstones, flysch, conglomerates). This simplified scheme captures key differences in the weathering behavior and mechanical response of the rock masses, which are known to influence the type and frequency of landslides. Fine-grained units generally produce cohesive soils prone to earth slides and flows, whereas coarse-grained units generate granular soils more susceptible to debris slides and debris flows. This categorization aligns with the "earth" and "debris" material types defined in the Cruden and Varnes (1996) landslide classification and is used throughout this study to interpret landslide patterns in relation to geological setting.



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173 **Table 1.** Simplified classification of the geological formations in the Emilia-Romagna region into
 174 eight litho-structural units, adapted from Berti et al. (2025). The classification is based on lithological
 175 composition and structural domain. Units 1 to 3 predominantly include fine-grained rocks, while units
 176 4 to 8 are composed mainly of coarse-grained lithologies.

Unit ID	Lithology	Domain	Structural position	Geological Age
1	Clays, silty clays, and marly clays	Padano-Adriatic	Outer Foredeep	Pliocene to Pleistocene
2	Marls and marly clays	Epiligurian	Wedge-top basins	Oligocene to Miocene
3	Clay shales, clay breccias, tectonized clays, olistostromes	Ligurian	Accretionary wedge	Cretaceous to Eocene
4	Massive rocks: basalts, serpentines, limestones, arenites	Ligurian, Epiligurian	Accretionary wedge Wedge-top basins	Cretaceous to Miocene
5	Flysch rocks made of rhythmic alternations of sandstones, limestones, pelites, and shales	Ligurian, Epiligurian	Accretionary wedge Wedge-top basins	Cretaceous to Eocene
6	Weakly cemented sandstones and conglomerates	Padano-Adriatic	Outer Foredeep	Pliocene to Pleistocene
7	Flysch rocks made of rhythmic alternations of sandstones and pelites	Tuscan-Umbrian	Inner Foredeep	Miocene
8	Weakly cemented sandstones with interbedded pelitic layers	Padano-Adriatic	Outer Foredeep	Pliocene to Pleistocene

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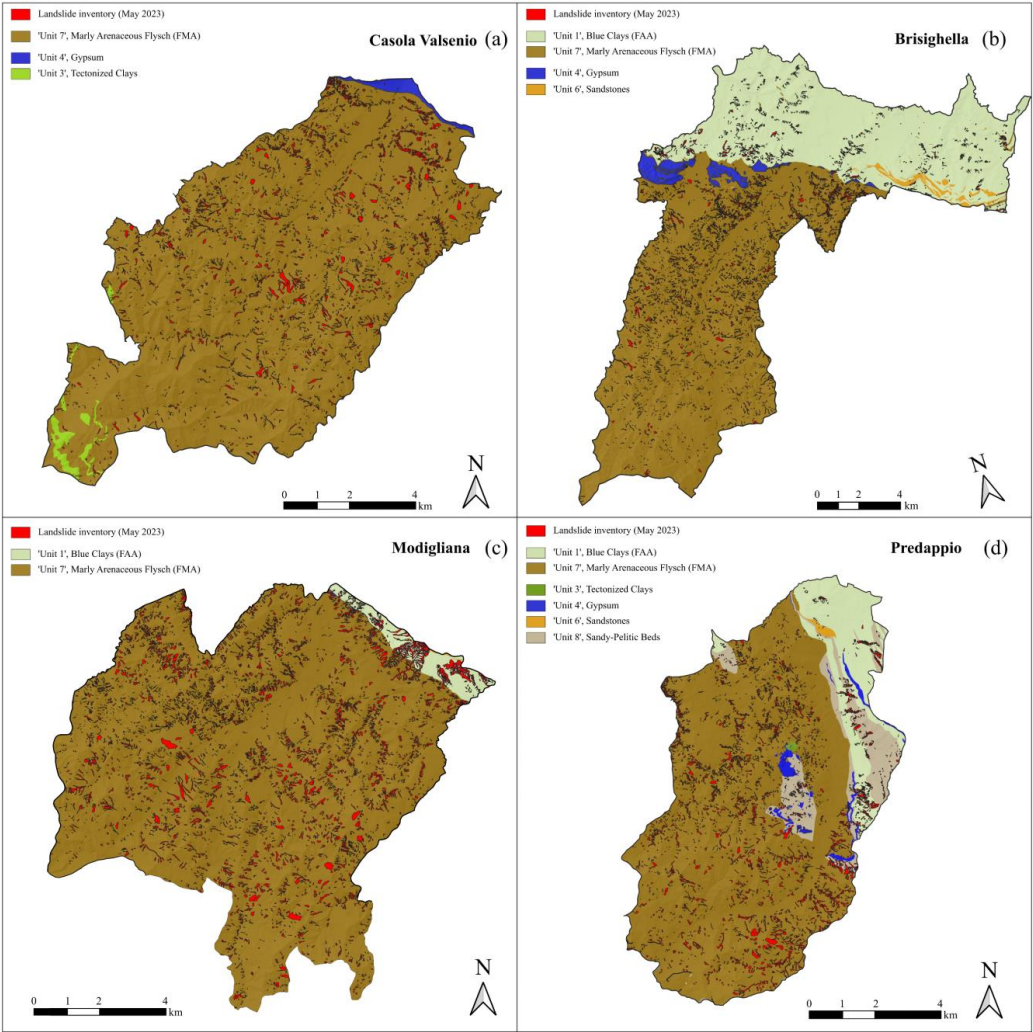


Figure 2. Schematic geological map of the municipalities of Casola Valsenio (a), Brisighella (b), Modigliana (c) and Predappio (d). FMA = Marnoso-Arenacea Formation, FAA = Blue Clays Formation.

Casola Valsenio

The municipality of Casola Valsenio is situated 30 km southwest of Imola (Fig. 1), covering an area of 84 km². It is predominantly known for its agricultural activities, including the cultivation of herbs, fruit orchards, and beekeeping. The underlying geology of the region is predominantly made up of the 'Unit 7' shows in Fig. 2a mainly composed by Marnoso-Arenacea Formation (FMA), a Tertiary flysch characterized by a finely interbedded sequence of marl and sandstone layers. This sequence was deposited in a deep marine foredeep basin during the Miocene epoch (Amy et al., 2006). The formation is characterized by a variable Arenites/Pelites (A/P) ratio, which reflects variations in



depositional environments within the basin. In Casola Valsenio, nearly the entire area features FMA with an A/P ratio ranging from 1/3 to 3/1. The northern sector is distinguished by predominantly sandstone-rich layers ($A/P > 3$) along with minor outcrops of Tectonized Clays (FPP) and Messinian Gypsum (FGY) formations. Before May 2023, the Geological Survey had documented 730 ancient, inactive landslides, primarily rock-block slides on cataclinal slopes that align with bedding planes. Following the May 2023 disaster, this figure dramatically increased to 5572 new landslides (Fig. 2a). Of these, debris slides (DS) and debris flows (DF) accounted for the vast majority, comprising 83% (4,543 events) and 14% (789 events) of the total, respectively. Typically, these were initiated by shallow failures of the thin soil layer (averaging 0.5-1 meters in depth) that overlies the Marnoso-Arenacea bedrock (Fig. 3b-c-d). In line with Berti et al. (2025), DF events were further classified into long-runout (DF1, 334 events) and limited-runout (DF2, 445 events), while DS events included 63 occurrences of limited-runout type (DS2, 1%). Additionally, 17 earth landslides were recorded, including earth flows (EF, 11 events; 0.2%) and earth slides (ES, 6 events; 0.1%). The area also witnessed several large rock-block slides on cataclinal slopes (Fig. 3a). Although these accounted for only 2% of the total landslides (110 events), their size was significantly greater than that of the debris slides and flows, resulting in extensive damage. The May 2023 disaster notably disrupted local infrastructure, blocking the main road and cutting off the southern region for weeks. Nearly 70 homes and 200 roadways were severely affected.

210 *Brisighella*

The municipality of Brisighella borders Casola Valsenio to the east (Fig. 1) and covers 194 km². The area is primarily composed of the 'Unit 7' composed by Marnoso-Arenacea Formation (FMA) (Fig. 2b). The northern part of the municipality, approximately 75 km² in size, consists mainly of the Blue Clays Formation (FAA) (Fine-grain rocks 'Unit 1' in Fig. 1b), which is bordered by a narrow strip of gypsum (Fig. 2b). Before the event, the Geological Survey had documented 2,100 landslides in the area, mostly ancient and inactive rock-block and complex slides. Following the disaster, an additional 6342 landslides were recorded.

Of these, 61% were debris slides (DS) (3,890 events), while debris flows (DF) accounted for 16% of the total (999 events), 6% of the with a long-runout (DF1) and 10% with limited-runout (DF2). Earth slides (ES) and earth flows (EF), typically associated with badlands in clayey terrains, represented 22% of the total (1,428 events). Lastly, rock-block slides accounted for 0.4% of the recorded events (25 occurrences). In the municipal territory, landslides blocked key access routes and damaged 66 homes and 261 roadways.

225 *Modigliana*

The municipality of Modigliana is located approximately 15 km to the east of Casola Valsenio (Fig. 1), and is one of the area most seriously affected by the event. The area spans 101 km² and is primarily composed of the Marnoso-Arenacea Formation, FMA ('Unit 7') with an A/P ratio ranging from 1/3 to 3/1 (Fig. 2c). A small portion of Blue Clays, FAA (Unit 1) outcrops in the northeast. Before the disaster, 795 landslides were documented, mostly complex and rock-block slides. After the May 2023 event, the number of landslides jumped to 6974, drastically reshaping the landscape. Of these, debris slides were the most common, accounting for 77% of the total (5357 events). Debris flows (DF) represented 19%, divided between long-runout (DF1, 573 events) and limited-runout (DF2, 733



234 events) types, while rock-block slides accounted for 1% of the total (69 events). Earth slides (ES) and
235 earth flows (EF) made up 3.5% (242 events). The disaster blocked the main access road for weeks,
236 affecting 51 homes and 155 roadways, and left the community to face long-term recovery challenges.

237

238 *Predappio*

239 The municipality of Predappio is situated around 30 km east of Casola Valsenio (Fig. 1) and spans an
240 area of 91 km². The geological framework of Predappio is mainly characterized by the Marnoso-
241 Arenacea Formation, FMA ('Unit 7') with an A/P ratio ranging from 1/3 to 3/1. The central part of
242 the municipality also features the Colombacci Formation ('Unit 8' in Fig. 1b), interspersed with
243 gypsum from the Gessoso-Solfifera Formation ('Unit 3'). Towards the northeast, the landscape
244 transitions to the Blue Clays formation ('Unit 1'), interbedded with deposits of the Colombacci
245 Formation and several occurrences of Pliocene sandstones (Borello Sandstones, 'Unit 7').

246 Before May 2023, 840 landslides had been recorded, primarily involving rock-block slides and debris
247 flows. Following the disaster, this number escalated to 6832. Of these, 80% were debris slides (5,476
248 events), while debris flows (DF) accounted for 11% (750) of the total including 419 long-runout
249 events (DF1) and 331 limited-runout events (DF2). Rock-block slides (RS), although only 0.5% of
250 the recorded events (31 occurrences), were notable for their significant size and potential impact.
251 Earth slides (ES) and earth flows (EF) represented 8% of the total (551 events).

252 The area experienced weeks of isolation due to several landslides blocking the main road, impacting
253 19 homes and damaging 138 roadways, again with long-lasting consequences.

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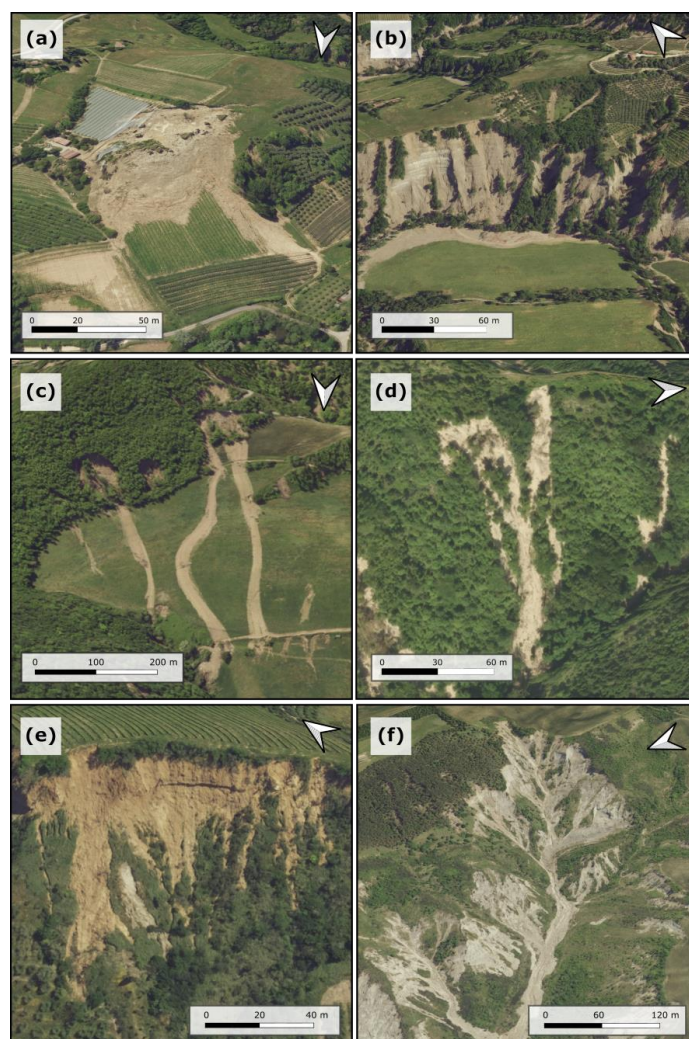


Figure 3. Examples of Landslides in Romagna from May 2023 event following Berti et al. 2025 classification.

a) Rockslide (RS) at Baccagnano, Brisighella, triggered during the night of May 2-3.
 (WGS84: 44.2069085, 11.7739792).

b) Debris slide (DS1) at Monte Albano, Casola Valsenio. (WGS84: 44.2356093, 11.6682287).

c) Debris flow with extended runout on gentle unforested slope (DF1) at Casola Valsenio.
 (WGS84: 44.20834, 11.64312).

d) Debris flow with limited runout on steep forested slope (DF2) at Modigliana.
 (WGS84: 44.15124, 11.75511).

e) Earth Slide (ES) at Brisighella. (WGS84: 44.23912, 11.78393).

f) Earth Flow (EF) at Predappio. (WGS84: 44.08947, 11.99627).

Images © Regione Emilia-Romagna, Geoportale 3D (<https://mappe.regione.emilia-romagna.it>)



3. Methods

3.1 Data

Deep-learning models for landslide mapping typically utilize a multi-layer input approach, integrating a combination of satellite or aerial imagery, topographical data, and land use information to capture the complex characteristics of landslides (Ghorbanzadeh et al., 2019, Meena et al., 2022). In our analysis, we considered the following seven layers (Fig. 4):

- L1) Post-event Sentinel-2 images: Four-bands (RGB+NIR) satellite images with a 10 m resolution, captured after the second rainfall event on May 23, 2023. They represent the first post-event Sentinel-2 images obtained with minimal cloud cover (Fig. 4b).
- L2) Pre-event Sentinel-2 images: Four-bands (RGB+NIR) satellite images with a 10 m resolution captured in May 2022, one year prior to the event. These images provide a similar vegetation state to that of May 2023, and were chosen to represent pre-event conditions due to the lack of cloud-free images in the two months preceding the event (Fig. 4a).
- L3) NDVI change Sentinel-2 map: Single-band raster generated by subtracting the pre-event NDVI map (Kriegler et al., 1969) derived from L2 from the post-event NDVI map created using L1 (Fig. 4c).
- L4) Post-event CGR images: Four bands (RGB+NIR) aerial photographs with a very high resolution of 0.2 m captured on May 23, 2023, shortly after the event. These images were committed by the Emilia-Romagna Region specifically for the disaster (Fig. 4f).
- L5) Pre-event AGEA images: Four bands (RGB+NIR) aerial photographs with a very high resolution of 0.2 m captured 3 years before the event (April to June 2020). These images were committed by the Agency for Agricultural Payments mainly for agricultural purposes (Fig. 4e).
- L6) NDVI change CGR map: Single-band raster generated by subtracting the pre-event NDVI map derived from L5 from the post-event NDVI map created using L4 (Fig. 4g).
- L7) Slope map: Single-band raster derived from the 5x5 m Digital Terrain Model (DTM) of the Emilia Romagna Region. The DTM, which is based on the 1:5000 scale Regional Technical Map, captures the pre-event slope morphology (Fig. 4d).

These layers were selected to simulate the diverse scenarios that might be encountered during an emergency, based on the available data. Table 2 details the potential combinations of the seven layers, with each combination corresponding to different scenarios explored in the study. In every case listed in Table 2, we assumed the availability of a slope map (L7), as numerous global DEMs such as SRTM (Farr et al., 2007; NASA SRTM, 2013), ASTER (NASA Earthdata Search, 2023), ALOS (Rosenqvist et al., 2007; ASF DAAC, 2023), and Copernicus DEM (ESA, 2019) are readily accessible for free.

Often, Sentinel-2 images (L1) might be the only data accessible following a large-scale disaster. The Sentinel-2 satellites, part of the European Copernicus program, revisit any point on Earth every 5 days and provide free 10 m resolution images in the visible and near-infrared spectrum, which are vital for various applications including disaster management (Wasowki et al., 2014; Yang et al., 2019; Ban et al., 2020). Numerous studies have utilized these open-access images to develop and evaluate automated methods for landslide mapping (Liu et al., 2021; Ghorbanzadeh et al., 2022; Lu et al.,



2023; Notti et al., 2023). Consequently, L1 can be considered the basic data layer available in the aftermath of an event (case 1 in Table 2).

Frequently, though not always, it is possible to access pre-event Sentinel-2 images (L2) and compute the NDVI change map (L3) to improve the identification of landslides triggered by an event. However, these resources may not be available if the disaster was preceded by extended periods of cloudy weather during which significant changes in vegetation state occurred. In our study, we utilized cloud-free Sentinel-2 images taken during the same period the previous year, but this approach isn't always feasible in regions with intensive agricultural activity, where the state of vegetation can vary markedly from one year to the next. When pre-event Sentinel-2 images are available, automated methods can be trained using L1, L2, and L3 (case 2).

In other cases, when financial resources permit, high-resolution images can be obtained soon after the disaster. These images may come from satellite systems such as WorldView-3 and 4 (Maxar Technologies, 0.31 m), GeoEye-1 (Maxar Technologies, 0.41 m), Pleiades 1A and 1B (Airbus, 0.5 m), IKONOS (Maxar Technologies, 0.82 m), or SkySat (Planet Labs, 0.9 m), or they can be captured through dedicated aerial campaigns. The use of high-resolution imagery greatly enhances the precision of landslide mapping as it allows for the detection of even the smallest features and the clear identification of subtle geomorphic signs of instability. When high-resolution post-event images are available, they can be used on their own (case 3), or in combination with Sentinel-2 images (case 4), facilitating the capture of both small-scale and large-scale landslide features.

Less commonly, high-resolution images suitable for analysis are available before the event. The utility of these images hinges on their timing relative to the disaster. Ideally, pre-event high-resolution images should capture the area immediately before the disaster to accurately pinpoint landslides triggered by the event and minimize false positives resulting from agricultural or natural vegetation changes. This requires appropriate satellite coverage of the area prior to the disaster. In our case, high-resolution aerial images were obtained three years before the event (L5); however, the general land cover remained unchanged, and no significant landslides had occurred in the area since then. Automated methods can be applied by integrating both pre- and post-event high-resolution images (case 5) or by also incorporating an NDVI change map derived from these datasets (case 6). A final case considers the combination of all available layers L1-7 (case 7).

Although the post-event (L4) and pre-event (L5) aerial images, as well as the corresponding NDVI change map (L6), were originally available at a very high spatial resolution of 0.2 m, their use at full resolution posed major computational issues. Specifically, attempts to process these layers at native resolution systematically resulted in out-of-memory (OOM) errors during model training and inference. This was due to the large size of the input tensors generated from 0.2 m imagery over wide areas, which exceeded the available GPU memory even when using optimized batch sizes and patching strategies.

To overcome this limitation and ensure stable execution of the deep learning pipeline, we resampled the L4, L5, and L6 layers to a coarser spatial resolution of 2 m. This down sampling allowed us to retain sufficient spatial detail for landslide identification, while significantly reducing the memory footprint of the input data. While the original 0.2 m resolution would have provided extremely fine detail, it exceeded the requirements and computational feasibility for a regional-scale deep learning analysis. In contrast, the 2 m resolution still captures the essential geomorphological features relevant



to landslide detection and represents a practical and effective compromise. This trade-off is consistent with the flexibility of deep-learning models, which can learn robust spatial representations from moderate-resolution inputs when properly trained and validated.

Table 2. Overview of the various input layer configurations used during the training processes. 'NDVI' refers to the 'ΔNDVI-CGR' Change map created using CGR imagery, while 'ΔNDVI-S2' refers to the NDVI Change map created using Sentinel-2 imagery.

Name	Model	Sentinel2 post [L1]	Sentinel2 pre [L2]	Sentinel2 ΔNDVI [L3]	CGR post [L4]	AGEA pre [L5]	CGR ΔNDVI [L6]	Slope [L7]
U1	U-Net	✓	✗	✗	✗	✗	✗	✓
S1	SegForm	✓	✗	✗	✗	✗	✗	✓
U2	U-Net	✓	✓	✓	✗	✗	✗	✓
S2	SegForm	✓	✓	✓	✗	✗	✗	✓
U3	U-Net	✗	✗	✗	✓	✗	✗	✓
S3	SegForm	✗	✗	✗	✓	✗	✗	✓
U4	U-Net	✓	✓	✓	✓	✗	✗	✓
S4	SegForm	✓	✓	✓	✓	✗	✗	✓
U5	U-Net	✗	✗	✗	✓	✓	✗	✓
S5	SegForm	✗	✗	✗	✓	✓	✗	✓
U6	U-Net	✗	✗	✗	✓	✓	✓	✓
S6	SegForm	✗	✗	✗	✓	✓	✓	✓
U7	U-Net	✓	✓	✓	✓	✓	✓	✓
S7	SegForm	✓	✓	✓	✓	✓	✓	✓

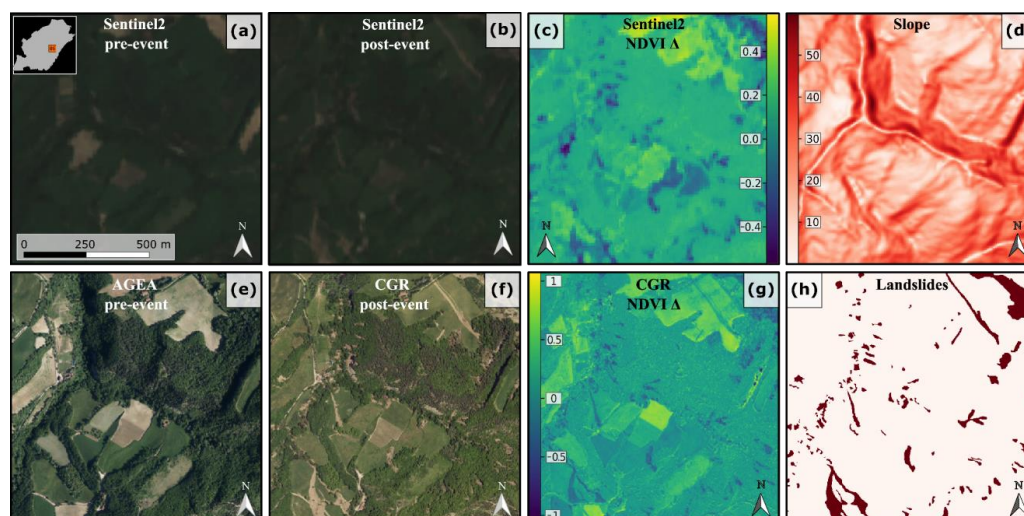


Figure 4. Example of the layers included in case 7, Tile No. 86. (a) Pre-event Sentinel-2 image (2022); (b) Post-event Sentinel-2 image; (c) NDVI change map derived from Sentinel-2 imagery; (d) Slope map; (e) AGEA imagery (2020), pre-event; (f) CGR aerial image (2023), post-event; (g) NDVI change map derived from high resolution imagery; (h) Binary raster showing manually mapped landslides (ground truth).

3.4 Deep Learning Semantic Segmentation Models

U-Net

U-Net is a convolutional neural network architecture originally designed for biomedical image segmentation by Ronneberger et al. (2015). It features an encoder-decoder structure where the encoder progressively reduces the spatial dimensions through convolutional and max-pooling layers, capturing contextual and high-level features. The decoder then up samples the features using transposed convolutions and concatenates them with corresponding encoder features through skip connections, which help retain spatial information and improve localization accuracy. This combination of context and localization has proven highly effective for segmentation tasks across various domains. U-Net's effectiveness in landslide mapping has been demonstrated in recent studies (Meena et al. 2021, 2022; Ghorbanzadeh et al. 2022, 2023; Nava et al. 2022).

In our study, we implemented the U-Net architecture manually in Python. The input shape of the images was defined according to the specific dimensions of the data used, ensuring the model could effectively process the input imagery. We set the number of scales to 2, which determines the number of downsampling steps in the encoder, and each scale included 2 convolutional layers, which helped capture intricate details and patterns within the data. The initial number of convolutional filters was set to 64. This parameter defines the number of filters applied in the first convolutional layer, which influences the model's ability to learn from the input data. As the network progresses through the layers, the number of filters increases, allowing the model to capture more complex features.

Our model was designed to output 2 classes: landslide and non-landslide. This binary classification approach enabled us to effectively differentiate between affected and unaffected areas in the imagery.



The model was compiled using the Adam optimizer (Kingma and Ba, 2017), a learning rate of 0.0001 and a Dice loss function for 300 epochs. The early stopping mechanism was introduced with a patience of 20 epochs to prevent overfitting, based on validation loss monitoring. All analyses were conducted using Python 3.8.18, TensorFlow 2.5.0, and CUDA 11.2.

SegFormer

SegFormer is a state-of-the-art architecture for semantic segmentation of multispectral landslide imagery (Xie et al., 2021). SegFormer, is designed to overcome limitations of previous vision Transformer models, such as their high computational cost, reliance on fixed input resolutions due to positional encodings, and inefficient multi-scale feature aggregation (Dosovitskiy et al., 2020). While the original SegFormer architecture was designed for 3-channel RGB images, the SegformerForSemanticSegmentation implementation from Hugging Face Transformers library (Wolf et al., 2020) allows for flexible input channel configuration. We leveraged this feature to adapt SegFormer to our multi channels imagery.

The SegFormer architecture consists of two main components (Xie et al., 2021):

1. Hierarchical Transformer Encoder: This encoder employs a novel hierarchical structure with progressively larger stride and patch sizes, allowing the model to capture both fine and coarse features efficiently. It uses an efficient self-attention mechanism without positional encoding and incorporates a Mix-FFN block that includes a depth-wise convolution to enhance local spatial information aggregation.
2. Lightweight All-MLP Decoder: Unlike complex decoder designs in previous models, SegFormer uses a simple multi-layer perceptron (MLP) decoder. This decoder fuses multi-level features from the encoder through a unified set of MLP layers, achieving per-pixel prediction without the need for complex operations like FPN or dilated convolutions.

SegFormer's design eliminates the need for positional encoding, allowing it to handle variable input resolutions without interpolation. This feature is particularly beneficial for our landslide imagery, which can vary in resolution and scale.

Our model was configured as the variant MiT-b0 (Xie et al., 2021) with the following parameters:

- Hidden sizes: [32, 64, 160, 256]
- Encoder depths: [2, 2, 2, 2]
- Decoder hidden size: 256

The model was trained using Cross-Entropy Loss as the criterion. We employed the Adam optimizer with a learning rate of 0.001 for 300 epochs. The early stopping mechanism was introduced with a patience of 20 epochs to prevent overfitting, based on validation loss monitoring. All analyses were conducted using Python 3.8.10, PyTorch (Paszke et al., 2019) 1.9.0, and CUDA 11.1.

3.5 Training-Testing Split

While conventional practice often dictates a 70/30 or 80/20 ratio for training and testing datasets (Hastie et al., 2009), we adopted a different approach to simulate real-world emergency mapping conditions. The model was exclusively trained on data from the Casola Valsenio municipality and



subsequently applied to three additional municipalities, Predappio, Modigliana, and Brisighella, for independent testing. This setup simulates a realistic emergency response scenario where; after mapping a relatively small area, the model is deployed to rapidly map other affected areas, relying solely on prior training.

To facilitate analysis, the train area was segmented into 1x1 km grid tiles, as illustrated in Fig. 5. Each 1 km² tile corresponds to 512x512 pixels in the CGR imagery when resampled at 2m resolution. Generally, a single grid tile in these areas encompasses over 50 landslides, providing a robust representation of landslide patterns.

For model training, we adopted a stratified random sampling approach. Out of the 97 total tiles in Casola Valsenio, 62% (60 tiles) were randomly allocated for training, 15% (15 tiles) for validation, and the remaining 23% (22 tiles) were reserved for initial testing and further validation (Fig. 5). This division was designed to maximize the training set while ensuring sufficiently large and reliable validation and test sets, given the need to apply the model to other municipalities.

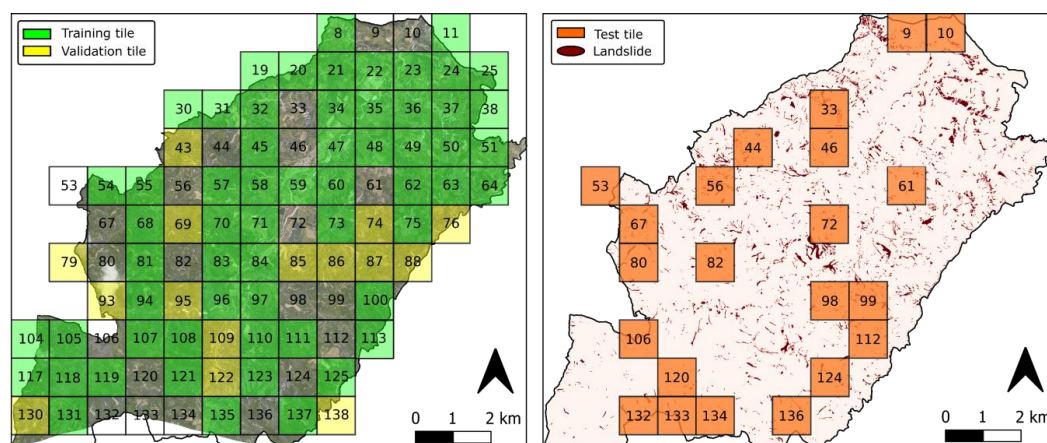


Figure 5. Training and test tiles used for models' construction in Casola Valsenio, illustrating the spatial split of the area.

3.6 Model Application

Following the training phase, we applied the models to the municipalities of Predappio, Modigliana, and Brisighella. Crucially, while ground truth (Fig. 1a) data was available for these municipalities, we deliberately chose not to use it for model training. Instead, this data was reserved exclusively for evaluating the performance of our models.

A significant constraint of the analysis arises from the predominance of the Marnoso-Arenacea Formation (FMA) in the training dataset from Casola Valsenio (Fig. 3a). To assess the models' ability to map landslides in varied geological contexts, we divided the Brisighella municipality into two distinct geological zones: one dominated by the Marnoso-Arenacea Formation (Brisighella FMA) and another characterized by the predominance of fine-grained rock from the Pliocene Blue Clays (Brisighella FAA) (Fig. 3d). This latter area is characterized by earth flows (EF) and earth slides (ES), which exhibit fundamentally different morphological, and soil cover features compared to landslides

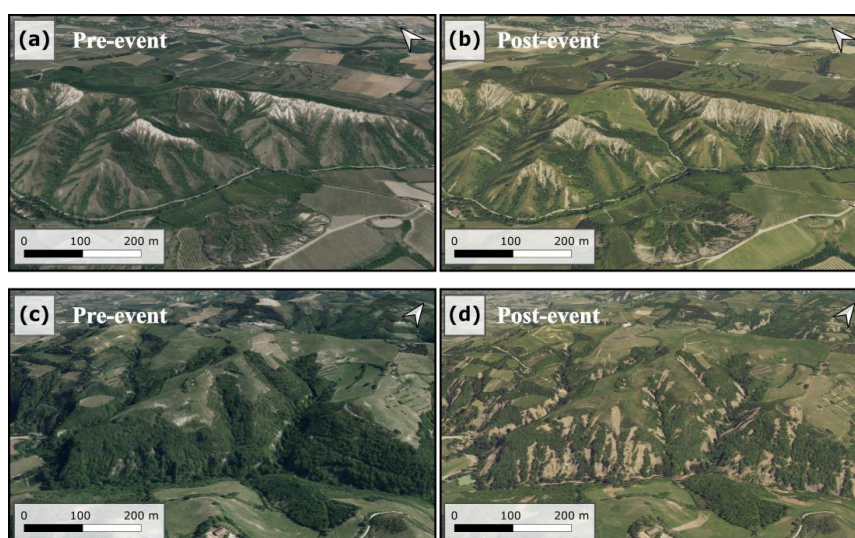


present in the Marnoso-Arenacea Formation. Figure 6 shows representative examples of these contrasting landslide types. The FAA zone, underlain by the Pliocene Blue Clays, is typified by slow-moving earth flows (EF) and earth slides (ES), which often occur within badland-like morphologies. These landslides commonly exhibit an elongated, teardrop-like form with a lobate toe, shaped by multiple episodes of reactivation over time. The vegetative cover is often sparse or degraded, as typically observed in clay-rich soils.

In contrast, landslides in the FMA zone, both in Brisighella and in the training area of Casola Valsenio, are dominated by debris slides (DS) and debris flows (DF). These failures typically originate from thin, unconsolidated colluvial layers that overlie the flysch bedrock, and are triggered by intense rainfall infiltration. The resulting movements are generally sudden, with well-defined scarp and runout zones, and affect areas with denser vegetation cover.

By explicitly separating these two lithological domains within Brisighella, we aimed to evaluate whether a model trained solely on FMA-type landslides could generalize to areas with fundamentally different geological and geomorphological settings. This setup also allowed us to assess the sensitivity of the model to lithology-driven differences in landslide type, geometry, and spatial pattern—an essential step toward developing mapping tools that are robust across variable terrain conditions.

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Figure 6. Comparison examples are shown between pre-event (a) and post-event (b) earth flow (EF) landslides in the Blue Clays Formation (FAA) and pre-event (c) and post-event (d) debris slide (DS) in the Marnoso-Arenacea Formation (FMA), both triggered by the May 2023 disaster in Brisighella. Images © Regione Emilia-Romagna, Geoportale 3D (<https://mappe.regione.emilia-romagna.it>)

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The final evaluation of model generalization was conducted on the entire set of grid tiles from the three remaining municipalities: 90 tiles from Predappio, 105 from Modigliana, and 196 from Brisighella. As such, although the core training-validation-testing split within Casola Valsenio



followed a 62/15/23 ratio, the effective test set includes not only the 22 held-out tiles from Casola Valsenio, but also 100% of the 413 tiles from the other three municipalities, a spatial hold-out test designed to assess cross-municipality generalization.

Table 3. Overview of the utilization of 1 km² tiles across the analyzed municipalities, illustrating how each area was segmented for model training, validation, testing and prediction purposes.

Area	Tiles [1 km ²]	Use
Casola Valsenio	97	60 train / 15 validation/ 22 test
Brisighella FMA	116	Prediction
Brisighella FAA	80	Prediction
Modigliana	105	Prediction
Predappio	90	Prediction

3.7 Evaluation Metrics and Expert Judgment

Metrics

To assess the performance of our landslide detection models, we employed two widely used metrics in semantic segmentation tasks: the F1 score and the Intersection over Union (IoU). These metrics were calculated on the entirety of the test images, rather than averaging scores across individual cells, to provide a comprehensive evaluation of the models' performance.

The F1 score, also known as the Dice coefficient, is the harmonic mean of precision and recall (Tharwat, 2020). It is particularly useful in imbalanced classification problems, such as landslide detection, where the positive class (landslide) is typically much smaller than the negative class (non-landslide). The F1 score is calculated as:

$$F1 = 2 \cdot \frac{Precision \cdot Recall}{Precision + Recall} \quad (1)$$

where Precision = TP / (TP + FP) and Recall = TP / (TP + FN), with TP, FP, and FN representing True Positives, False Positives, and False Negatives, respectively.

The Intersection over Union (IoU), also known as the Jaccard index, is another common metric for assessing segmentation accuracy (Rezatofighi et al., 2019). It measures the overlap between the predicted segmentation mask and the ground truth mask. The IoU is calculated as:

$$IoU = \frac{TP}{TP + FP + FN} \quad (2)$$

Both metrics range from 0 to 1, with 1 indicating perfect prediction. By using these metrics on the entire test images, we can evaluate how well our models perform in detecting and delineating landslides across varied landscapes and geological contexts.

Expert Judgment

In automated landslide mapping, a higher algorithmic score does not necessarily correspond to a higher quality map (Zhang et al., 2018; Isensee et al., 2021). This issue is particularly relevant in



516 automated landslide mapping, where the practical utility of the map often depends more on its
 517 interpretability and coherence with geomorphological reality than on its statistical performance alone.

518 For example, a model might overpredict large landslide areas to optimize pixel-wise metrics like IoU,
 519 while neglecting smaller or fragmented features that are critical for field validation and emergency
 520 response. Conversely, models may introduce noise or irregular boundaries that artificially increase
 521 recall but reduce the practical readability of the map.

522 While this issue is typically addressed by selecting evaluation metrics that emphasize specific
 523 components of the confusion matrix (e.g., recall over precision), in emergency scenarios not all
 524 components carry the same operational weight. For instance, in crisis management, a higher recall
 525 may be preferred to avoid missing active landslides, even if this comes at the cost of more false
 526 positives. In contrast, in densely urbanized areas or during resource-constrained interventions,
 527 precision may become more important to avoid unnecessary alarms and misallocation of efforts.

528 To address this challenge, three of the authors in this study provided their expert judgment on the
 529 quality of the automatically generated landslide maps. This evaluation was conducted without prior
 530 knowledge of the scores obtained by these maps or the training input layers used to generate them.

531 The assessment focused on seven common error types frequently encountered in automated landslide
 532 mapping:

533 E1 – False positives in agricultural fields, where cultivated land is mistakenly classified as landslide;

534 E2 – False positives along roads and in urban areas, where the absence of vegetation may be
 535 misinterpreted as landslide activity;

536 E3 – False positives along riverbeds, often due to confusion with recent sediment scouring or bedload
 537 deposits;

538 E4 – False negatives in areas with known landslides;

539 E5 – Over segmentation, referring to the fragmentation of a single landslide into multiple small
 540 polygons;

541 E6 – Inaccurate delineation of landslide boundaries;

542 E7 – Omission errors in shadowed areas, such as steep, shaded slopes or regions obscured by cloud
 543 cover.

544 For each error type, the experts assigned a score of 0 (limited error), 1 (moderate error), or 2 (relevant
 545 error) to indicate the severity of the issue. These scores were then summed and normalized to calculate
 546 an Expert-based Performance Index (EPI):

547
$$EPI = \frac{14 - \sum_{i=1}^{i=7} E_i}{14}$$

548 The index ranges from 0 to 1, with 1 representing the highest mapping performance. This evaluation
 549 was independently carried out by each of the three experts for all 16 cases included in the analysis
 550 (see Table 2). The experts assessed the model cases without prior knowledge of the specific
 551 configurations, thereby minimizing potential bias.

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554 **4. Results**

555 4.1 Model results and performance

556 The results of the analysis are presented in Table 4. The table details the F1 and IoU scores obtained
557 by comparing the automated landslide maps produced by two machine learning models (U=U-Net;
558 S=SegFormer) against the manually created maps. The rows list the seven different combinations of
559 input layers, numbered from 1 to 7, to represent a progressively richer dataset (refer to Table 1). The
560 models were initially trained using data from Casola Valsenio and subsequently applied to four
561 additional areas (Predappio, Modigliana, Brisighella FMA, and Brisighella FAA). Overall, this
562 combination of 7 input datasets, 4 areas, and 2 models resulted in 56 automated landslide maps.

563 In the tested configurations, the models U6, U7, and S7 achieved the highest F1 and IoU scores. These
564 models incorporate nearly all (U6) or all (U7-S7) of the seven input layers, and clearly benefit from
565 the high-resolution pre- and post-event images. However, the performance gains from utilizing all
566 available layers are relatively modest. Notably, model S2, which employs only Sentinel-2 images and
567 Sentinel-2 NDVI, achieves results close to those of the top-performing models, with differences in
568 F1 and IoU scores frequently around a margin of ~0.01. Similar observations apply to models U4 and
569 S4, which rely on Sentinel-2 products and post-event high-resolution images. This suggests that while
570 increasing input information enhances model performance, even more streamlined configurations can
571 still yield robust and competitive mapping outcomes. The good performance of U4 and S4 models
572 highlights the effectiveness of integrating high-resolution imagery alongside Sentinel-2-derived
573 indices, even when fewer spectral inputs are available.

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Table 4. Quality of the automatic mapping, based on F1 and IoU scores, achieved by U-Net and SegFormer algorithms with different combinations of input layers, detailed in Table 2.

Models' Name	Casola Valsenio		Predappio		Modigliana		Brisighella FMA		Brisighella FAA	
	F1	IoU	F1	IoU	F1	IoU	F1	IoU	F1	IoU
U1	0.56	0.39	0.42	0.27	0.48	0.31	0.50	0.33	0.32	0.19
S1	0.47	0.31	0.46	0.30	0.45	0.29	0.46	0.30	0.34	0.21
U2	0.64	0.47	0.52	0.35	0.55	0.38	0.57	0.40	0.38	0.24
S2	0.62	0.45	0.54	0.37	0.57	0.40	0.57	0.40	0.47	0.31
U3	0.70	0.54	0.39	0.24	0.45	0.29	0.55	0.38	0.29	0.17
S3	0.54	0.37	0.44	0.28	0.44	0.29	0.49	0.32	0.32	0.19
U4	0.72	0.56	0.55	0.38	0.58	0.41	0.62	0.45	0.47	0.31
S4	0.68	0.52	0.59	0.42	0.60	0.43	0.62	0.45	0.49	0.33
U5	0.71	0.55	0.53	0.36	0.56	0.39	0.62	0.45	0.50	0.34
S5	0.65	0.48	0.56	0.39	0.56	0.39	0.60	0.43	0.43	0.28
U6	0.72	0.56	0.56	0.38	0.59	0.42	0.63	0.46	0.53	0.36
S6	0.65	0.48	0.53	0.36	0.54	0.37	0.58	0.41	0.41	0.26
U7	0.73	0.57	0.56	0.39	0.60	0.42	0.63	0.46	0.53	0.36
S7	0.66	0.49	0.60	0.43	0.60	0.43	0.61	0.44	0.52	0.35

Figure 7 shows the F1-score trends across all seven input layer combinations (cases 1 to 7) for the five evaluated areas. In both models, a marked increase in performance is observed between case 1 and case 2, particularly in Casola Valsenio, Modigliana, and Brisighella FAA. This highlights the importance of including NDVI change information alongside Sentinel-2 imagery to reach better segmentation performance. From case 2 onward, the models tend to stabilize, with U-Net showing a more pronounced benefit from additional input layers, especially in cases 4 to 7. SegFormer, by contrast, reaches near-plateau performance earlier and exhibits smaller gains as layers are added.

Performance across the different municipalities reveals further differences between the two models. U-Net shows a clear separation in F1-scores depending on the area: it performs best in Casola (reaching 0.73 in case 7), moderately well in Brisighella FMA, Modigliana, and Predappio, and significantly worse in Brisighella FAA (with 0.29 in case 3 and 0.53 in case 7). This trend reflects U-Net's sensitivity to geological conditions.

SegFormer, on the other hand, provides more uniform scores across areas, although generally lower than U-Net in Casola. Its performance ranges between 0.47–0.68 in Casola and remains around 0.55–0.60 in Predappio, Modigliana, and Brisighella FMA. Brisighella FAA again represents a challenge,



with F1-scores fluctuating and even slightly declining from case 2 onward, suggesting that additional input layers do not consistently improve results in this more complex setting.

Overall, both models identify Brisighella FAA as the most difficult area, with the lowest F1-scores recorded. However, while U-Net appears to benefit more from richer input combinations, SegFormer demonstrates greater generalization ability across areas with different lithological and morphological characteristics.

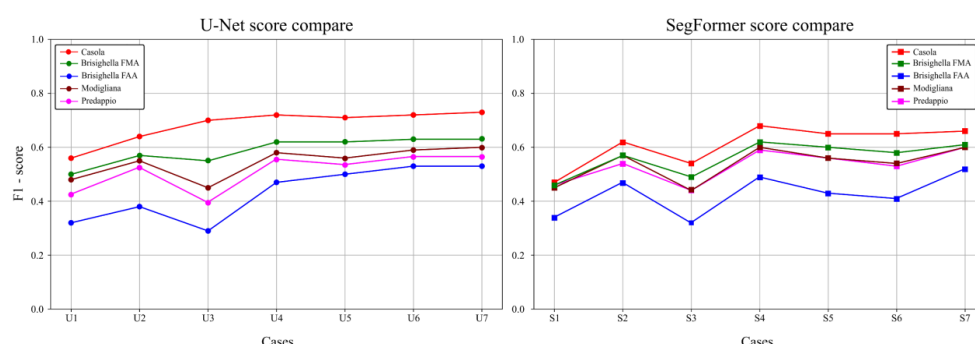


Figure 7. F1-score comparison across all seven input layer combinations (cases 1–7) for each municipality, as obtained by the U-Net and SegFormer models.

As a complement to the quantitative evaluation, Figures 8 and 9 provide illustrative examples of model performance in real-world conditions. These visual comparisons, focusing on the best-performing configurations (U-Net case 7 and SegFormer case 7), highlight both successful detections and common failure modes observed across the test municipalities.

Figure 8 displays cases where both models performed well, accurately mapping landslides in areas with widespread vegetation removal, typical of debris slides (DS) and debris flows (DF). These landslide types, which together accounted for about 81% of all events in May 2023, were consistently recognized, especially in severely impacted zones where spectral and texture changes were prominent. Figure 9, by contrast, presents a selection of cases where both models encountered notable difficulties, leading to different types of misclassifications:

- Casola Valsenio: Despite being the training area, the models occasionally misinterpreted the spatial extent of landslide polygons. In some cases, large landslides are fragmented into several smaller segments, resulting in False Negatives and underestimation of the actual affected area. This issue is likely due to the complex shape and internal heterogeneity of the landslide footprints.
- Modigliana: Both U7 and S7 show a tendency to produce False Positives in cultivated fields. These areas are subject to seasonal changes due to agricultural activities such as plowing and vegetation regrowth, which often produce surface variations that resemble the spectral signatures of recent landslides. In particular, disturbed or freshly tilled soil patches in post-event imagery may be incorrectly classified as landslides.
- Brisighella FMA: In this area, False Positives were frequently mapped along riverbanks, likely caused by localized color differences between the pre- and post-event images. These



640 differences may be associated with flood-related sediment deposition or changes in water
 641 level and channel morphology following localized inundations. Additionally, some False
 642 Positives were recorded on recently constructed buildings, especially where bare roofs or
 643 fresh concrete surfaces produced strong spectral contrasts. These elements, not present in
 644 the pre-event imagery, appear as “new” disturbed areas, misleading the models.

- 645 ▪ Brisighella FAA: This area poses the most significant challenge for both models.
 646 Dominated by badlands developed in Blue Clay formations—geologically and
 647 morphologically distinct from the Marnoso-Arenacea Formation of the training area—
 648 Brisighella FAA yielded high rates of False Positives. The stark lithological difference
 649 appears to have limited the models’ ability to generalize to this terrain, underlining the
 650 importance of including lithological variability in training datasets.
- 651 ▪ Predappio: In this area, errors are primarily due to landslides occurring in shadowed
 652 regions of the post-event imagery. The lower contrast and altered color profiles in these
 653 zones reduced the models’ ability to correctly detect landslide features, leading to partial
 654 detection or complete omission (False Negatives). These results suggest that terrain
 655 shading, common in steep valleys, remains a limitation for optical-image-based mapping.

656 Together, these examples illustrate both the strengths and the current limitations of the models
 657 when applied in real-world scenarios. While debris-type landslides in well-illuminated,
 658 vegetated settings were reliably mapped, generalization to areas with different lithologies,
 659 land use, or illumination conditions remains an area for improvement.

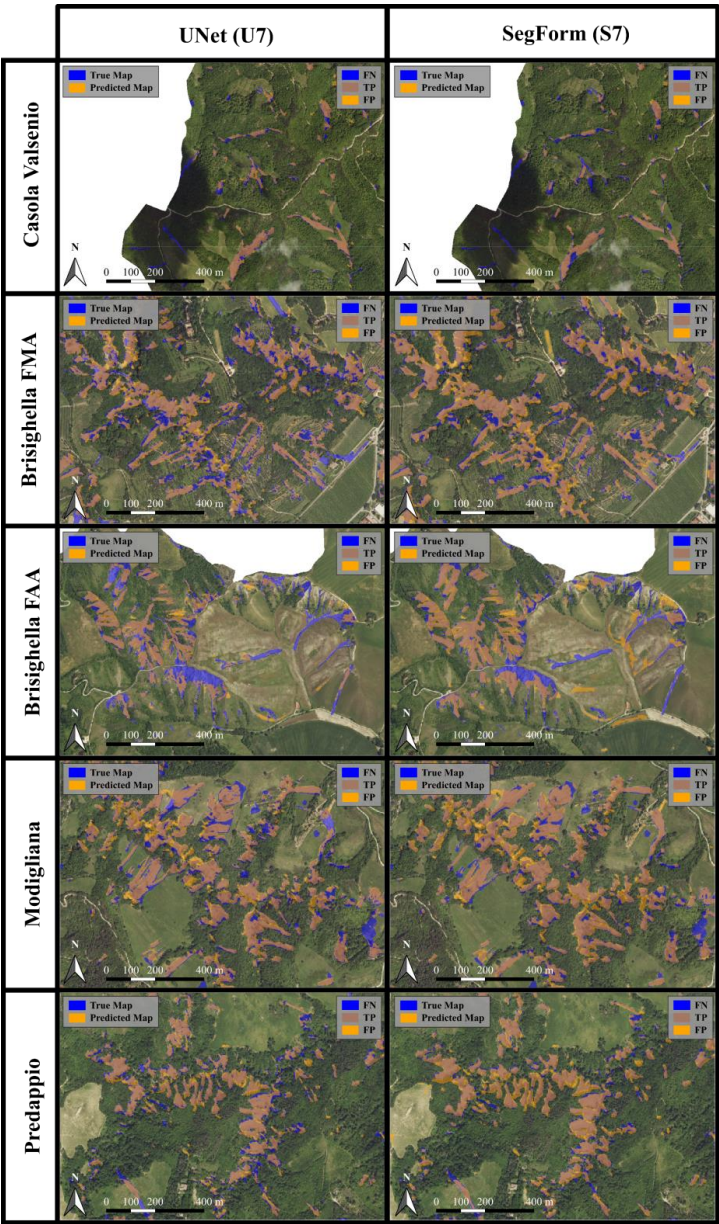


Figure 8. Comparative analysis of landslide mapping results using U7 and S7 models across different municipalities. The figure illustrates the spatial distribution and extent of landslides as identified by the U7 (UNet-based) and S7 (SegFormer-based) models in Casola Valsenio, Predappio, Modigliana, Brisighella FMA, and Brisighella FAA. Both models were trained using pre- and post-event imagery, slope data, and NDVI from CGR. The comparison highlights the similarities and differences in landslide detection capabilities between the two models, particularly in areas dominated by Debris Slides (DS).

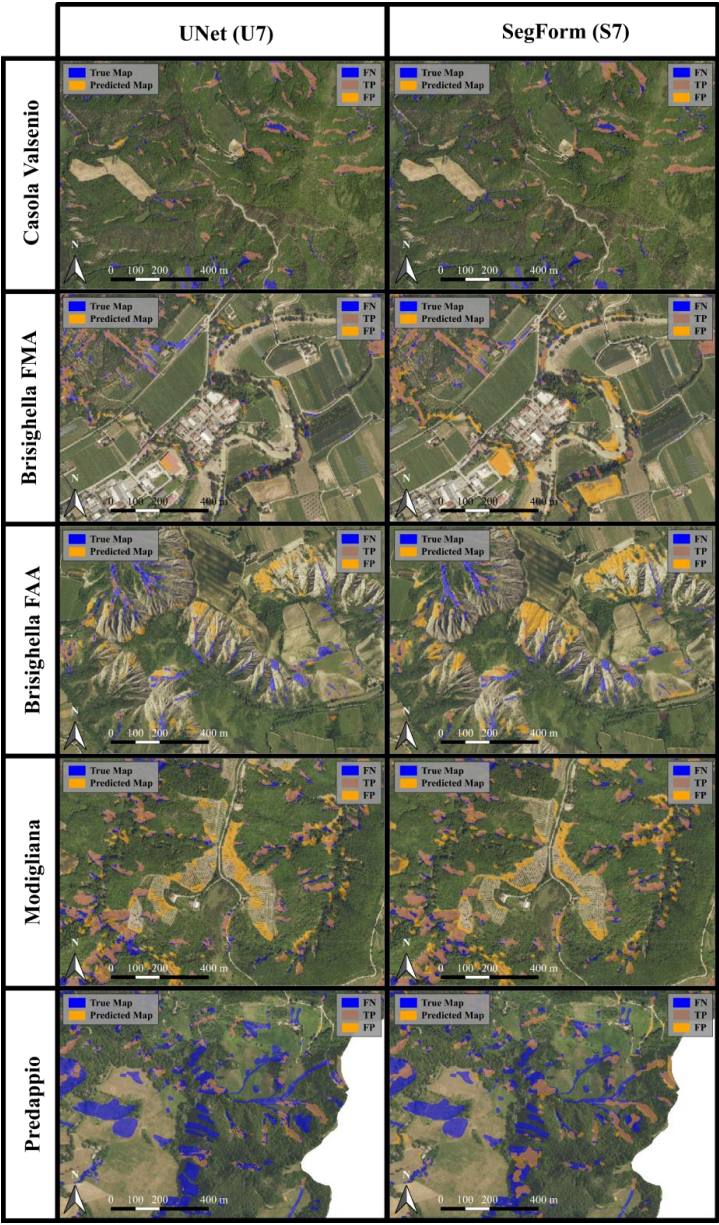


Figure 9. Representative examples of challenging scenarios for landslide detection using U7 and S7 models across municipalities. This figure illustrates common situations where both models encounter difficulties in accurately mapping landslides, highlighting the limitations and potential areas for improvement in the machine learning approach. The examples are drawn from Casola Valsenio, Predappio, Modigliana, Brisighella FMA, and Brisighella FAA, showcasing the diverse geological and environmental conditions that can impact model performance.



4.2 Expert judgement

Figure 10 presents a comparison between the F1-scores of all model cases and the EPI (Expert Performance Index) values derived from expert assessments (see Section 3.7). The EPI values shown in the charts reflect evaluations from three experts, who assessed the automated landslide maps produced by the two models and assigned weights to seven common detection errors.

In general, the F1-scores align well with the EPI values obtained from expert judgement. Although the F1-scores exhibit limited variability, the trend of improved performance with increasing model complexity is also shown by the expert-based evaluations. Notably, several performance fluctuations are similarly interpreted: model U3 is clearly recognized as underperforming, while model U7 is consistently considered one of the best by both quantitative scores and expert opinion.

Some notable discrepancies emerge for certain models that were rated as the best by a single expert but not by F1-scores or the other experts. For instance, model S5 received the highest score from Expert 1, model U2 from Expert 2, and model U4 from Expert 3, despite not being ranked similarly by others. These differences reflect the specific areas each expert emphasized during evaluation, as well as the inherent subjectivity of expert judgment. However, the trend of the mean EPI (dashed lines in Fig.10), which helps reduce individual biases, closely follows that of the F1-score.

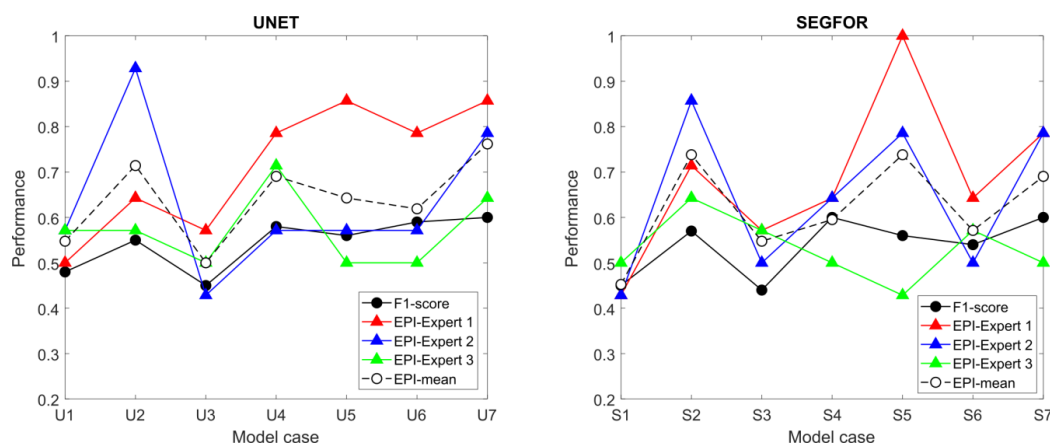


Figure 10. Comparison of model performance based on F1-scores and expert-based evaluations from three independent experts.

5. Discussion

One of the key observations from this study concerns the differing sensitivity of U-Net and SegFormer to input data combinations, particularly in the context of emergency response mapping. U-Net's performance varies significantly depending on the layers used, while SegFormer yields more consistent results across different configurations. This contrast is likely rooted in their architectural differences. U-Net, designed for pixel-wise segmentation, relies heavily on the spatial and spectral characteristics of the input data, making it more sensitive to the quality and availability of supplementary layers such as pre-event imagery or NDVI change maps. SegFormer, being a



transformer-based model, is better equipped to capture global context, thus reducing dependency on specific input combinations and improving robustness under complex geological conditions.

This architectural contrast becomes even more evident when models are transferred across different geological domains. As shown in the results (Figure 7), U-Net exhibits a marked decline in performance when moving from areas geologically similar to the training site (e.g., Predappio, Modigliana) to Brisighella FAA, where lithological conditions differ significantly. This suggests that U-Net has limited generalization capacity beyond the Marnoso-Arenacea Formation (FMA), and struggles in fine-grained, clay-rich terrains like those found in the FAA domain. SegFormer, although generally more stable, also underperforms in Brisighella FAA, highlighting the broader challenge of transferring models across heterogeneous geological settings.

To further investigate the role of geology in model generalization, an additional analysis was conducted to assess the impact of including lithology among the input layers. For instance, comparing Case S3 (RGB + slope) with the same configuration plus a lithology layer did not yield significant improvements in Casola, Predappio, Modigliana, or Brisighella MA. These areas are predominantly characterized by the Marnoso-Arenacea Formation (“Unit 7”), which is already well represented in the training dataset (Figure 2a), meaning the lithology layer provides redundant information. However, when the same model was applied to Brisighella FAA, where lithology is dominated by Blue Clays (“Unit 1”), it failed to detect any landslides, resulting in a very low F1-score of 0.02. This highlights a key limitation: simply adding lithology as a static input layer is insufficient for ensuring generalization. If the training dataset is not lithologically balanced, the model may reinforce existing biases, associating landslide occurrence primarily with “Unit 7” and failing to detect events in “Unit 1”.

This finding underscores a fundamental point: effectively incorporating lithological information into AI-based landslide mapping requires more than providing it as an input feature. The training dataset itself must include landslide polygons from a representative range of lithological settings. When the model was explicitly trained and tested on Brisighella FAA using the same configuration (Case S3: RGB + slope), the F1-score increased from 0.32 to 0.41, confirming that geological consistency in the training data substantially improves the model’s ability to detect landslides in previously underrepresented domains.

Although the F1-scores achieved in this study may appear modest compared to some machine learning benchmarks (<https://huggingface.co/papers/trending>), they represent a substantial accomplishment in the context of landslide mapping. Unlike conventional image classification tasks, landslide mapping requires the precise delineation of complex and irregular shapes across highly variable terrain. Our models aim to produce a realistic representation of landslide extents, including subtle transitions and intricate morphologies, rather than simplified or generalized outputs. This level of accuracy is difficult to attain, yet essential for practical applications in emergency response and hazard assessment.

In this regard, the potential of these automated approaches in emergency scenarios could be considerable. The ability to rapidly generate landslide maps of comparable quality to expert-drawn products could greatly accelerate initial response efforts. By reducing the need for time-consuming manual digitization, these methods could save weeks of work, an especially critical advantage during crisis events (Berti et al. 2025). The automatically generated maps may serve as a robust initial



product, allowing practitioners to focus on refinement and validation rather than starting from scratch, thereby enabling the delivery of a high-quality final map in a fraction of the time required for full manual mapping.

A further consideration is the use of automated landslide maps for identifying damage and at-risk structures. Following the May 2023 disaster, the Emergency Commission determined that all buildings located within landslide boundaries or within a 20-metre buffer were to be considered at risk and potentially subject to relocation. To evaluate the usefulness of our automated mapping products for this task, we compared the buildings identified automatically with those derived from manual mapping. Figure 11 presents the confusion matrix for the 16 modelled cases. The F1-scores indicate that “U2”, “U4”, “S6”, and “U7” are the most accurate CNN outputs, successfully identifying on average 528 out of 654 buildings at risk (~0.78 F1-score). These results demonstrate the models’ ability to estimate the spatial extent of the phenomenon, even when using Sentinel-2 imagery at 10 m resolution (U2). Nevertheless, all models produced both false positives (incorrectly flagged buildings) and false negatives (overlooked buildings). Consequently, manual verification remains essential when high accuracy is required, and misclassifications cannot be tolerated.

Even a small number of misclassified structures may have serious consequences when inhabited buildings are involved. This raises a key question: how can model performance be further improved to support more effective and precise emergency response efforts in the future? Several avenues could be explored:

1. Higher-resolution input data: As discussed in Section 3.1, high-resolution aerial imagery (20 cm/pixel) and NDVI change maps were down sampled to 2 m/pixel to reduce computational load. Using the original resolution, while computationally expensive, could yield improved detail and model performance (Wang et al., 2022).
2. Improved tiling strategies: The current approach relies on fixed 496×496 px tiles for training, validation, and testing (Section 3.5). This can split landslide polygons at tile boundaries and reduce spatial context. As noted by Abrahams et al. (2024), employing sliding windows or full-scene processing could mitigate these issues and enhance segmentation quality.
3. Data augmentation: No augmentation techniques were used in this study. Simple operations such as flipping, rotation, or brightness adjustment, shown by Prakash et al. (2021) could improve CNN robustness reducing overfitting and improve generalizability.

Finally, although current approaches enable rapid mapping of landslides, their ability to assess structural exposure remains limited. Buildings located just outside mapped polygons may still be at significant risk, which a uniform 20 m buffer may fail to capture adequately. A more advanced solution is proposed by BGC Engineering (2010), who incorporate topographic features and estimated landslide runout directions to dynamically assess exposure. In this framework, risk is evaluated based on terrain morphology, flow paths, and downslope connectivity, offering a more realistic representation of potential impacts compared to distance-based buffers. Such models could represent a major step forward in automated damage assessment, complementing segmentation outputs with geomorphologically informed prioritization strategies.



Input Layers		U-NET	SegForm																								
Case 1	<ul style="list-style-type: none">• Post-event S2• Slope	<div>U1</div> <div>F1: 0.53</div> <div>Predicted value</div> <table><tr><td rowspan="2">Outside</td><td>TN</td><td>FP</td></tr><tr><td>1822</td><td>781</td></tr><tr><td rowspan="2">Inside</td><td>FN</td><td>TP</td></tr><tr><td>139</td><td>515</td></tr></table> <div>Actual value</div> <div>Outside Inside</div>	Outside	TN	FP	1822	781	Inside	FN	TP	139	515	<div>S1</div> <div>F1: 0.67</div> <table><tr><td>Outside</td><td>TN</td><td>FP</td></tr><tr><td></td><td>2221</td><td>382</td></tr><tr><td>Inside</td><td>FN</td><td>TP</td></tr><tr><td></td><td>137</td><td>517</td></tr></table> <div>Outside Inside</div>	Outside	TN	FP		2221	382	Inside	FN	TP		137	517		
		Outside		TN	FP																						
1822	781																										
Inside	FN	TP																									
	139	515																									
Outside	TN	FP																									
	2221	382																									
Inside	FN	TP																									
	137	517																									
Case 2	<ul style="list-style-type: none">• Post-event S2• Pre-event S2• Δ NDVI-S2• Slope	<div>U2</div> <div>F1: 0.78</div> <table><tr><td>Outside</td><td>TN</td><td>FP</td></tr><tr><td></td><td>2455</td><td>148</td></tr><tr><td>Inside</td><td>FN</td><td>TP</td></tr><tr><td></td><td>144</td><td>510</td></tr></table> <div>Outside Inside</div>	Outside	TN	FP		2455	148	Inside	FN	TP		144	510	<div>S2</div> <div>F1: 0.77</div> <table><tr><td>Outside</td><td>TN</td><td>FP</td></tr><tr><td></td><td>2421</td><td>182</td></tr><tr><td>Inside</td><td>FN</td><td>TP</td></tr><tr><td></td><td>134</td><td>520</td></tr></table> <div>Outside Inside</div>	Outside	TN	FP		2421	182	Inside	FN	TP		134	520
		Outside	TN	FP																							
	2455	148																									
Inside	FN	TP																									
	144	510																									
Outside	TN	FP																									
	2421	182																									
Inside	FN	TP																									
	134	520																									
Case 3	<ul style="list-style-type: none">• Post-event CGR• Slope	<div>U3</div> <div>F1: 0.72</div> <table><tr><td>Outside</td><td>TN</td><td>FP</td></tr><tr><td></td><td>2497</td><td>106</td></tr><tr><td>Inside</td><td>FN</td><td>TP</td></tr><tr><td></td><td>230</td><td>424</td></tr></table> <div>Outside Inside</div>	Outside	TN	FP		2497	106	Inside	FN	TP		230	424	<div>S3</div> <div>F1: 0.68</div> <table><tr><td>Outside</td><td>TN</td><td>FP</td></tr><tr><td></td><td>2261</td><td>342</td></tr><tr><td>Inside</td><td>FN</td><td>TP</td></tr><tr><td></td><td>137</td><td>517</td></tr></table> <div>Outside Inside</div>	Outside	TN	FP		2261	342	Inside	FN	TP		137	517
		Outside	TN	FP																							
	2497	106																									
Inside	FN	TP																									
	230	424																									
Outside	TN	FP																									
	2261	342																									
Inside	FN	TP																									
	137	517																									
Case 4	<ul style="list-style-type: none">• Post-event S2• Pre-event S2• Δ NDVI S2• Post-event CGR• Slope	<div>U4</div> <div>F1: 0.78</div> <table><tr><td>Outside</td><td>TN</td><td>FP</td></tr><tr><td></td><td>2415</td><td>188</td></tr><tr><td>Inside</td><td>FN</td><td>TP</td></tr><tr><td></td><td>116</td><td>538</td></tr></table> <div>Outside Inside</div>	Outside	TN	FP		2415	188	Inside	FN	TP		116	538	<div>S4</div> <div>F1: 0.73</div> <table><tr><td>Outside</td><td>TN</td><td>FP</td></tr><tr><td></td><td>2285</td><td>318</td></tr><tr><td>Inside</td><td>FN</td><td>TP</td></tr><tr><td></td><td>90</td><td>564</td></tr></table> <div>Outside Inside</div>	Outside	TN	FP		2285	318	Inside	FN	TP		90	564
		Outside	TN	FP																							
	2415	188																									
Inside	FN	TP																									
	116	538																									
Outside	TN	FP																									
	2285	318																									
Inside	FN	TP																									
	90	564																									
Case 5	<ul style="list-style-type: none">• Post-event CGR• Pre-event AGEA• Slope	<div>U5</div> <div>F1: 0.76</div> <table><tr><td>Outside</td><td>TN</td><td>FP</td></tr><tr><td></td><td>2387</td><td>216</td></tr><tr><td>Inside</td><td>FN</td><td>TP</td></tr><tr><td></td><td>117</td><td>537</td></tr></table> <div>Outside Inside</div>	Outside	TN	FP		2387	216	Inside	FN	TP		117	537	<div>S5</div> <div>F1: 0.75</div> <table><tr><td>Outside</td><td>TN</td><td>FP</td></tr><tr><td></td><td>2377</td><td>226</td></tr><tr><td>Inside</td><td>FN</td><td>TP</td></tr><tr><td></td><td>123</td><td>531</td></tr></table> <div>Outside Inside</div>	Outside	TN	FP		2377	226	Inside	FN	TP		123	531
		Outside	TN	FP																							
	2387	216																									
Inside	FN	TP																									
	117	537																									
Outside	TN	FP																									
	2377	226																									
Inside	FN	TP																									
	123	531																									
Case 6	<ul style="list-style-type: none">• Post-event CGR• Pre-event AGEA• Δ NDVI CGR• Slope	<div>U6</div> <div>F1: 0.73</div> <table><tr><td>Outside</td><td>TN</td><td>FP</td></tr><tr><td></td><td>2290</td><td>313</td></tr><tr><td>Inside</td><td>FN</td><td>TP</td></tr><tr><td></td><td>97</td><td>557</td></tr></table> <div>Outside Inside</div>	Outside	TN	FP		2290	313	Inside	FN	TP		97	557	<div>S6</div> <div>F1: 0.78</div> <table><tr><td>Outside</td><td>TN</td><td>FP</td></tr><tr><td></td><td>2436</td><td>167</td></tr><tr><td>Inside</td><td>FN</td><td>TP</td></tr><tr><td></td><td>125</td><td>529</td></tr></table> <div>Outside Inside</div>	Outside	TN	FP		2436	167	Inside	FN	TP		125	529
		Outside	TN	FP																							
	2290	313																									
Inside	FN	TP																									
	97	557																									
Outside	TN	FP																									
	2436	167																									
Inside	FN	TP																									
	125	529																									
Case 7	<ul style="list-style-type: none">• Post-event S2• Pre-event S2• Δ NDVI-S2• Post-event CGR• Pre-event AGEA• Δ NDVI-CGR• Slope	<div>U7</div> <div>F1: 0.79</div> <table><tr><td>Outside</td><td>TP</td><td>FP</td></tr><tr><td></td><td>2436</td><td>167</td></tr><tr><td>Inside</td><td>FN</td><td>TN</td></tr><tr><td></td><td>120</td><td>534</td></tr></table> <div>Outside Inside</div>	Outside	TP	FP		2436	167	Inside	FN	TN		120	534	<div>S7</div> <div>F1: 0.75</div> <table><tr><td>Outside</td><td>TP</td><td>FP</td></tr><tr><td></td><td>2282</td><td>321</td></tr><tr><td>Inside</td><td>FN</td><td>TN</td></tr><tr><td></td><td>71</td><td>583</td></tr></table> <div>Outside Inside</div>	Outside	TP	FP		2282	321	Inside	FN	TN		71	583
		Outside	TP	FP																							
	2436	167																									
Inside	FN	TN																									
	120	534																									
Outside	TP	FP																									
	2282	321																									
Inside	FN	TN																									
	71	583																									

786

787 **Figure 11.** Comparison of buildings identified as at risk (within landslide boundaries or within a 20-
788 meter buffer) by automated mapping methods and manual mapping (ground truth), showing the
789 confusion matrices for all five cases evaluated.



790

791 6. Conclusion

792 In this work, we investigated the potential of automated landslide mapping to support rapid
 793 emergency response following the extreme meteorological events of May 2023 in Emilia-Romagna,
 794 Italy. Using a deep learning approach, we trained and tested two semantic segmentation models, U-
 795 Net and SegFormer, on high-resolution aerial, Sentinel-2 imagery, NDVI change maps, and slope
 796 data. The training was conducted solely in the municipality of Casola Valsenio, while model
 797 performance was assessed on three additional municipalities (Predappio, Modigliana, Brisighella),
 798 chosen for their geological settings and significant landslide occurrence.

799 The main findings of the study can be summarized as follows:

- 800 ▪ **Training performance:** In the training area of Casola Valsenio, both models achieved high
 801 performance when provided with rich input combinations. U-Net reached F1 scores around
 802 0.70, particularly when both high-resolution imagery and NDVI change maps were included
 803 (Figure 7). These results confirm that, when trained on spatially detailed and spectrally rich
 804 datasets, CNN-based models can closely replicate manually mapped landslide boundaries,
 805 including small and morphologically complex features.
- 806 ▪ **Generalization to test areas:** When applied to unseen areas, both U-Net and SegFormer
 807 produced comparable F1 scores across Predappio, Modigliana, and Brisighella MA. While
 808 slight variations were observed, SegFormer occasionally performing better in certain
 809 configurations, there was no systematic advantage indicating stronger generalization. In
 810 Brisighella FAA, both models experienced a clear performance drop, highlighting the
 811 challenge of transferring models to geologically distinct terrains (Table 4, Figure 7). These
 812 results suggest that generalization is primarily constrained by lithological dissimilarity, rather
 813 than by architectural differences between CNN and transformer-based models.
- 814 ▪ **Model comparison:** Despite their architectural differences, the two models achieved similar
 815 performance across all test areas. This convergence likely reflects the limited variability and
 816 discriminative power of the available input layers, suggesting that both architectures were
 817 able to extract the most relevant features from the data (Figure 7). Improving the variety and
 818 informativeness of the input layers could unlock further improvements, especially in
 819 geologically heterogeneous regions.
- 820 ▪ **F1 vs. Expert Index:** The Expert Performance Index (EPI) was largely consistent with the
 821 quantitative F1 scores, reinforcing the validity of standard segmentation metrics for assessing
 822 model quality (Figure 10). Experts noted that the best-performing models also produced
 823 outputs that were more coherent from a geomorphological perspective, but they also identified
 824 recurring issues such as over segmentation, misclassification in agricultural areas, and false
 825 positives along riverbanks. These errors, while systematic, are easily corrected during manual
 826 review, making the model highly valuable as a starting point for rapid emergency mapping.
- 827 ▪ **Lithology influence:** The inclusion of lithology as an additional input layer had limited effect
 828 when the model was applied to areas already represented in the training dataset. However, in
 829 Brisighella FAA, where lithologies were almost absent from the training set, both models
 830 performed poorly unless explicitly retrained. This highlights that effective integration of



lithological information into AI-based mapping requires more than adding it as an input feature: the training dataset must include landslide polygons across a representative range of lithological contexts. When the model was retrained on Brisighella FAA, the F1-score increased, confirming that geological consistency in the training data significantly improves generalization performance.

- **Exposure and damage detection:** The automated maps successfully identified a majority of the buildings located within or near landslide polygons, reaching an F1 score of approximately 0.78 for building-at-risk detection (Figure 11). However, both false positives (e.g., buildings wrongly marked as at risk) and false negatives (e.g., undetected at-risk buildings) were frequent. This limits the reliability of automated exposure assessments, particularly in densely populated or infrastructure-rich areas. More advanced exposure analysis methods that account for slope direction, runout potential, and terrain morphology (e.g., BGC Engineering, 2010) are needed to complement segmentation-based approaches.

Overall, automated landslide mapping offers substantial advantages in emergency scenarios. While not a substitute for manual efforts, it provides a reliable baseline that can significantly accelerate the production of actionable maps. With further refinement, such as full-resolution imagery (Wang et al., 2022), improved tiling strategies (Abrahams et al., 2024), and data augmentation (Prakash et al., 2021), these tools can become critical assets in disaster management workflows. Given the increasing scale and frequency of climate-driven disasters, integrating AI-based approaches into civil protection protocols could enhance both the speed and consistency of emergency responses. Deep learning models, despite current limitations, represent a scalable solution to the growing demand for rapid, accurate, and spatially detailed hazard information. By automatically addressing the most evident features, these models reduce the burden on experts, allowing them to focus on verifying complex or ambiguous regions. Nonetheless, manual validation remains essential, as AI outputs alone cannot yet ensure sufficient reliability in sensitive contexts.

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Disclosures and declarations

Declaration of AI and AI-assisted technologies in the writing process

During the preparation of this work the authors used ChatGPT 4 (chat.openai.com) to enhance the grammar and syntax, as well as to refine the sentence structure. All the content is original, and no concepts, ideas, or interpretations were produced by this tool. After using this tool, the authors



867 reviewed and edited the content as needed and take full responsibility for the content of the
 868 publication.

869

870 **Code and data availability**

871 The dataset used in this study is openly available at <https://doi.org/10.5281/zenodo.13742643>.

872 The code developed for the analyses is available from the authors upon reasonable request.

873

874 **Author contribution**

875 NDS performed the data analysis, developed the scripts, and wrote the manuscript.

876 GC contributed to the interpretation of the results and supervised the work.

877 DE contributed to the methodology and algorithmic framework.

878 ELP contributed to the methodology and algorithmic framework.

879 AC supervised the interpretation of the results.

880 MB conceived the study, supervised the analysis and manuscript preparation, and provided critical

881 review and guidance throughout the research process.

882

883 **Competing interests**

884 The authors declare they have no financial interests.

885

886 **References**

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