

Rapid Landslide Mapping During the 2023 Emilia-Romagna Disaster: Assessing Automated Approaches with Limited Training Data

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The catastrophic rainfall events of May 2023 in the Emilia-Romagna region, Italy, triggered more than 80,000 landslides, as documented in a publicly available inventory (<https://doi.org/10.5281/zenodo.13742643>), and placed extraordinary demands on emergency response systems. One of the most critical emergency tasks was landslide mapping, which was carried out manually and required substantial time and effort. This study investigates the potential of automated landslide mapping to support rapid disaster response by evaluating two deep learning models, U-Net and SegFormer, under realistic emergency constraints, including limited training data.

The models were trained using data from one affected municipality, Casola Valsenio (84 km²), and tested on three additional municipalities, Predappio (91 km²), Modigliana (101 km²), and Brisighella (194 km²), characterized by different geological settings. To represent a range of operational scenarios, we tested seven combinations of input data, from post-event Sentinel-2 imagery alone to the integration of high-resolution aerial imagery, vegetation change maps, and slope data. Both models achieved comparable segmentation performance, with SegFormer showing greater robustness to variations in input data and geological conditions, while U-Net was more sensitive but occasionally more accurate when richer inputs were available. Both models successfully identified landslides, but showed limitations in shadowed areas, cultivated fields, and geologically distinct terrains. A major limitation emerged in the Brisighella area, where poor generalization was associated with the dominance of Blue Clay formations and the limited lithological diversity of the training data. These results highlight the importance of geologically balanced datasets for improving model transferability.

Overall, the study confirms the operational value of automated landslide mapping as a first-pass tool for emergency response. Although manual revision remains necessary, both models produced reliable baseline maps that can support validation and help prioritize interventions in time-critical situations, offering a scalable and time-efficient approach to the growing need for rapid and spatially detailed hazard information.

1. Introduction

Rapid and accurate landslide mapping over large areas is a challenging task, particularly in the context of regional-scale rainfall or earthquake events (Iverson et al. 2015; Casagli et al. 2016; Robinson et al., 2017; Holbling et al. 2017; Mondini et al., 2021). This task is crucial for implementing timely and effective response measures, assessing the extent of impact, and planning for recovery and mitigation efforts. The complexities of rapid landslide mapping vary depending on several factors including the extent of the affected area, the availability of cloud-free imagery, the precision needed for the mappings, and the difficulties of detecting landslides due to shadows, vegetation cover, or weak geomorphic evidences (Mondini et al. 2019; Amatya et al., 2023). Typically, these complexities are addressed by expert geologists through the visual interpretation of satellite or aerial imagery (Guzzetti et al., 2012; Scaioni et al., 2014; Ferrario and Livio 2023). These experts are skilled at identifying subtle variations in the terrain and can integrate complex contextual knowledge, including field reports, personal accounts, and specific geological information, to create highly accurate landslide maps. However, manual mapping is both time-consuming and inherently subjective, which are significant limitations during an emergency (Novellino et al. 2024).

Before the widespread adoption of deep learning, automated landslide mapping approaches evolved through several methodological stages. Early methods relied mainly on pixel-based techniques, such as thresholding of spectral indices or terrain parameters, which were simple to implement but highly sensitive to noise and illumination conditions. These approaches were later complemented by object-based image analysis (OBIA), which segments images into meaningful objects and classifies them based on spectral, spatial, and contextual features (Blaschke 2010; Martha et al. 2010). OBIA-based methods represented a substantial improvement by incorporating shape, texture, and neighborhood information, and were widely applied to landslide detection from high-resolution imagery. However, both pixel- and object-based approaches typically depend on manually designed rules or handcrafted features, limiting their transferability across different events, sensors, and geological settings.

More recently, machine-learning and deep-learning approaches have progressively replaced rule-based systems by enabling data-driven feature extraction and end-to-end learning. Automated landslide recognition techniques using convolutional neural networks (CNNs) have emerged as promising alternatives to manual mapping (Sameen and Pradhan 2019; Chen et al. 2018; Ye et al. 2019; Tang et al. 2022; Ji 2020). Numerous studies have demonstrated their effectiveness in real-world post-disaster scenarios. For example, Meena et al. (2021) applied a deep learning approach for rapid landslide mapping in India following extreme monsoon rainfall, while Prakash et al. (2021) introduced a generalized CNN framework designed to improve applicability across different geographic contexts. Prakash and Manconi (2021) further demonstrated the operational value of CNN-based mapping by rapidly delineating landslides triggered by severe storm events. Beyond individual case studies, recent research has focused on transferability, reproducibility, and systematic comparison of methods. Prakash et al. (2020) explicitly compared deep-learning models with traditional machine-learning approaches for landslide mapping from Earth observation data, highlighting the advantages of deep architectures in complex environments. In parallel, benchmark datasets such as Landslide4Sense (Ghorbanzadeh et al. 2022), HR-GLDD (Meena et al. 2023), and the CAS Landslide Dataset (Xu et al. 2024) have been released to promote standardized evaluation across multiple events, sensors, and geographic regions. These efforts have significantly advanced

the field by enabling more robust inter-comparisons and by revealing persistent challenges related to class imbalance, spatial heterogeneity, and generalization across geological domains.

In parallel with CNN developments, transformer-based architectures have recently been introduced in landslide mapping to better capture long-range spatial dependencies and multi-scale contextual information. Unlike CNNs, which primarily exploit local convolutional kernels, transformers rely on self-attention mechanisms that allow each pixel (or patch) to attend to a broader spatial context. This property is particularly relevant for landslide detection, where the geomorphological setting, slope connectivity, and spatial coherence of failures extend beyond local neighborhoods. Hybrid CNN–Transformer models and pure transformer-based architectures have shown promising results in improving robustness and generalization across heterogeneous landscapes (Chen et al. 2023; Wu et al. 2024). These characteristics are especially attractive for rapid response applications, where models must be deployed with limited training data and applied to large, spatially diverse areas under operational constraints.

Despite these advances, a substantial gap remains between methodological developments and their practical deployment during emergency response. Most deep-learning studies rely on relatively large and well-balanced training datasets, often adopting train-to-validation ratios such as 80:20 or 70:30 (Meena et al. 2023). In contrast, during an ongoing crisis only a small fraction of the affected area is typically mapped and annotated in the first hours to days, making extensive training datasets unrealistic. This limitation is often exacerbated by delays in acquiring cloud-free post-event imagery, inconsistencies in pre-event reference data, and the operational need to deploy models rapidly with minimal manual intervention. Furthermore, early training samples are frequently not representative of the full variability of landslide processes and environmental conditions, including different landslide types, lithologies, land-cover patterns, and illumination or shadow effects. As a result, models trained on limited and potentially biased early annotations may generalize poorly when applied to the broader affected region.

In this study, we utilize the landslide inventory from the May 2023 crisis in the Emilia-Romagna region of Italy (Berti et al. 2025) to assess the effectiveness and limitations of automated mapping algorithms in a practical setting. We specifically focus on the performance of the U-Net and SegFormer algorithms in identifying landslides triggered by the event, despite the challenges of very limited training data. The study also considers the variety of landslide types and geological conditions prevalent in the area, which impacted the disaster response. The main goal of this research is to determine whether advanced deep-learning methods can effectively replace manual mapping in managing large-scale landslide disasters, thus bridging the gap between academic research and real-world application in emergency management.

2. The Romagna May 2023 disaster

2.1 Overview of the disaster

In May 2023, the Emilia-Romagna region of Italy was hit by unprecedented meteorological events, as detailed by Berti et al. (2025) and Foraci et al. (2023). Two major rainfall episodes, from May 1-3 and May 16-17, led to severe floods and landslides primarily affecting the eastern part of the region (Fig. 1a). The first event (May 1-3) was driven by a stationary low-pressure system which delivered over 200 mm in 48 hours. The second event (May 16-17), featured a similar low-pressure system

enhanced by bora winds and moist air from the Mediterranean, resulting in rainfall exceeding 250 mm in the same duration. The cumulative impact of these episodes resulted in more than 500 mm of rainfall over two weeks, nearly half the region's annual climatic precipitation (Fig. 1a). The return period for these combined rainfall events was estimated to exceed 1000 years (Brath et al. 2023), highlighting the unprecedented nature of this occurrence. The two cyclonic events led to extensive flooding, submerging 540 km² of land, and triggered more than 80,000 landslides across approximately 1000 km² (Fig. 1b). This widespread devastation impacted thousands of buildings, roads, and other infrastructure, with the estimated damages amounting to approximately 9 billion euros.

Immediately after the disaster, we initiated the development of a detailed inventory map of the landslides caused by the event. This map was manually created using high-resolution aerial imagery taken only one week following the second rainfall event. The aerial images, which include four bands (RGB+NIR) and have a resolution of 0.2 m, allowed for the accurate identification and classification of the landslides. The data derived from this mapping effort are described in Berti et al. (2025) and are available for public download in shapefile format (DOI: 10.5281/zenodo.13742643, Pizziolo et al., 2024).

2.2 Motivation of the work

The visual identification of the landslides from May 2023 proved to be a challenging and time-intensive process. It took six months for twelve expert geologists to identify and delineate all the landslides within the 1000 km² area affected by the disaster. Despite these challenges, this approach was deemed the best available method at the time to meet the Civil Protection agency's specific requirements for accuracy, precision, and oversight during the emergency response coordination.

As this manual mapping progressed, we explored various automated mapping techniques to speed up the process and reduce the workload. Initial trials using NDVI methods and U-Net in a specific sample area produced promising results, demonstrating that deep-learning models can effectively replicate manual mapping efforts (Berti et al., 2026). However, these preliminary analyses were conducted in a test area (the Casola Valsenio municipality, Fig. 1), which covers only about 10% of the total area impacted by the event. Consequently, several critical questions remain unanswered: i) Can automated models trained on a small area be effectively used for rapid mapping of larger regions? ii) How do models trained in specific geological conditions perform across different geological settings? iii) Could more sophisticated deep-learning methods surpass U-Net in improving outcomes?

Addressing these questions is crucial to determine whether emerging automated methods can enhance the efficiency of landslide mapping in future crisis scenarios, thereby potentially reducing critical response times. In this study, we simulate landslide mapping during a crisis, taking into account the limited data typically available and the accessible information under such conditions. This simulation is grounded in our firsthand experience of the real emergency during the May 2023 event, which provided us with valuable insights into the practical challenges and data limitations encountered during such crises.

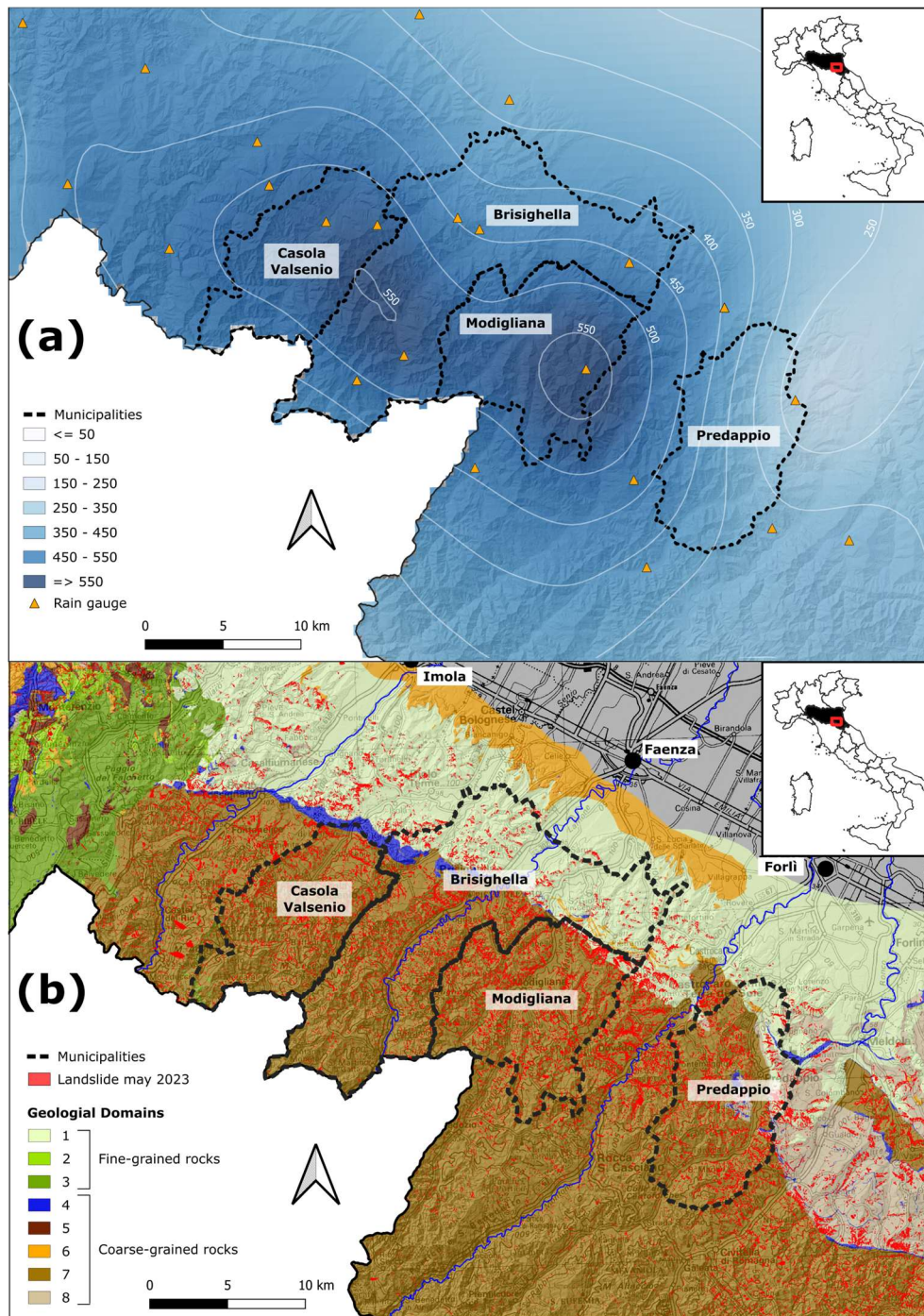


Figure 1. Overview of the Effects of the Extreme Rainfall Event in Romagna, May 2023.

a) Cumulative precipitation from 1 to 17 May 2023 derived from the regional rain-gauge network operated by Regione Emilia-Romagna. **b)** Manually mapped landslide distribution across the study areas, categorized by lithological units (Units 1-8) following the classification proposed by Berti et al. (2025; Table 2).

173 3. Study area

174 *Casola Valsenio*

175 The municipality of Casola Valsenio is situated 30 km southwest of Imola (Fig. 1), covering an area
176 of 84 km². It is predominantly known for its agricultural activities, including the cultivation of herbs,
177 fruit orchards, and beekeeping. The underlying geology of the region is predominantly made up of
178 the ‘Unit 7’ mainly composed of the Marnoso-Arenacea Formation (FMA) (Unit 7 in the lithological
179 classification of Berti et al., 2025; Table 2), a Tertiary flysch characterized by a finely interbedded
180 sequence of marl and sandstone layers. This sequence was deposited in a deep marine foredeep basin
181 during the Miocene epoch (Amy et al., 2006). The formation is characterized by a variable
182 Arenites/Pelites (A/P) ratio, which reflects variations in depositional environments within the basin.
183 In Casola Valsenio, nearly the entire area features FMA with an A/P ratio ranging from 1/3 to 3/1.
184 The northern sector is distinguished by predominantly sandstone-rich layers (A/P > 3) along with
185 minor outcrops of Tectonized Clays (FPP) and Messinian Gypsum (FGY) formations. Before May
186 2023, the Geological Survey had documented 730 ancient, inactive landslides, primarily rock-block
187 slides on cataclinal slopes that align with bedding planes. Following the May 2023 disaster, this figure
188 dramatically increased to 5572 new landslides (Fig. 2a). Landslide types and counts reported here are
189 derived from the official RER2023 inventory (Berti et al., 2025), which is openly available as a vector
190 dataset on Zenodo (Pizziolo et al., 2024; <https://doi.org/10.5281/zenodo.13742643>). The post-event
191 landslide inventory in Casola Valsenio is dominated by debris slides and debris flows, mainly
192 triggered by shallow failures of the thin soil cover overlying the Marnoso-Arenacea bedrock (Fig.
193 3b–d; Table 1). Both long- and limited-runout debris flows are observed, reflecting differences in
194 slope morphology and vegetation. Earth movements are rare, whereas rock-block slides, although
195 less frequent, are locally significant due to their larger size and damage potential (Fig. 3a). The May
196 2023 disaster notably disrupted local infrastructure, blocking the main road and cutting off the
197 southern region for weeks. Nearly 70 homes and 200 roadways were severely affected.

198

199 *Brisighella*

200 The municipality of Brisighella borders Casola Valsenio to the east (Fig. 1) and covers 194 km². The
201 area is primarily composed of the ‘Unit 7’ composed by Marnoso-Arenacea Formation (FMA) (Fig.
202 2b). The northern part of the municipality, approximately 75 km² in size, consists mainly of the Blue
203 Clays Formation (FAA) (Fine-grain rocks ‘Unit 1’ in Fig. 1b), which is bordered by a narrow strip of
204 gypsum (Fig. 2b). Before the event, the Geological Survey had documented 2,100 landslides in the
205 area, mostly ancient and inactive rock-block and complex slides. Following the disaster, an additional
206 6342 landslides were recorded. Landslide processes vary markedly with lithology in this area. Debris
207 slides prevail in areas underlain by the Marnoso-Arenacea Formation, while earth slides and earth
208 flows are more common in the Blue Clays sector, where badland morphologies are widespread (Table
209 1). Debris flows occur with variable runout, and rock-block slides are relatively rare but locally
210 impactful.

211

212 *Modigliana*

213 The municipality of Modigliana is located approximately 15 km to the east of Casola Valsenio (Fig.
214 1), and is one of the area most seriously affected by the event. The area spans 101 km² and is primarily

215 composed of the Marnoso-Arenacea Formation, FMA ('Unit 7') with an A/P ratio ranging from 1/3
 216 to 3/1 (Fig. 2c). A small portion of Blue Clays, FAA (Unit 1) outcrops in the northeast. Before the
 217 disaster, 795 landslides were documented, mostly complex and rock-block slides. After the May 2023
 218 event, the number of landslides jumped to 6974, drastically reshaping the landscape. The landslide
 219 inventory in Modigliana is largely dominated by debris slides, with debris flows representing a
 220 secondary but relevant process (Table 1). Both long- and limited-runout debris flows contributed to
 221 slope and channel instability, whereas rock-block slides and earth movements occur only sporadically.
 222 This pattern reflects widespread shallow instability within the Marnoso-Arenacea Formation.

223

224 *Predappio*

225 The municipality of Predappio is situated around 30 km east of Casola Valsenio (Fig. 1) and spans an
 226 area of 91 km². The geological framework of Predappio is mainly characterized by the Marnoso-
 227 Arenacea Formation, FMA ('Unit 7') with an A/P ratio ranging from 1/3 to 3/1. The central part of
 228 the municipality also features the Colombacci Formation ('Unit 8' in Fig. 1b), interspersed with
 229 gypsum from the Gessoso-Solfifera Formation ('Unit 3'). Towards the northeast, the landscape
 230 transitions to the Blue Clays formation ('Unit 1'), interbedded with deposits of the Colombacci
 231 Formation and several occurrences of Pliocene sandstones (Borello Sandstones, 'Unit 7'). Debris
 232 slides, here, represent the most common landslide type, accompanied by numerous debris flows with
 233 variable runout behavior (Table 1). Rock-block slides are infrequent but potentially severe, while
 234 earth slides and earth flows are more common than in purely sandstone-dominated areas, reflecting
 235 the local presence of clay-rich lithologies. The area experienced weeks of isolation due to several
 236 landslides blocking the main road, impacting 19 homes and damaging 138 roadways, again with long-
 237 lasting consequences.

238

239 **Table 1.** Summary of the landslides mapped in the four study municipalities after the May 2023
 240 Emilia-Romagna event, including total counts and breakdown by landslide type. DS = debris slides;
 241 DF = debris flows; DF1 = long-runout debris flows; DF2 = limited-runout debris flows; ES = earth
 242 slides; EF = earth flows; RS = rock-block slides. Data are from the RER2023 landslide inventory
 243 (Berti et al., 2025) and its open-access release (Pizziolo et al., 2024).

244

Municipality	Tot. Landslides	Debris Slides (DS)	Debris Flows (DF)	Earth Slides (ES)	Earth Flows (EF)	Rock Slides (RS)
Casola Valsenio	5572	4543	789	6	11	110
Brisighella	6342	3890	999	835	593	25
Modigliana	6974	5357	1306	152	90	69
Predappio	6832	5476	750	426	117	31

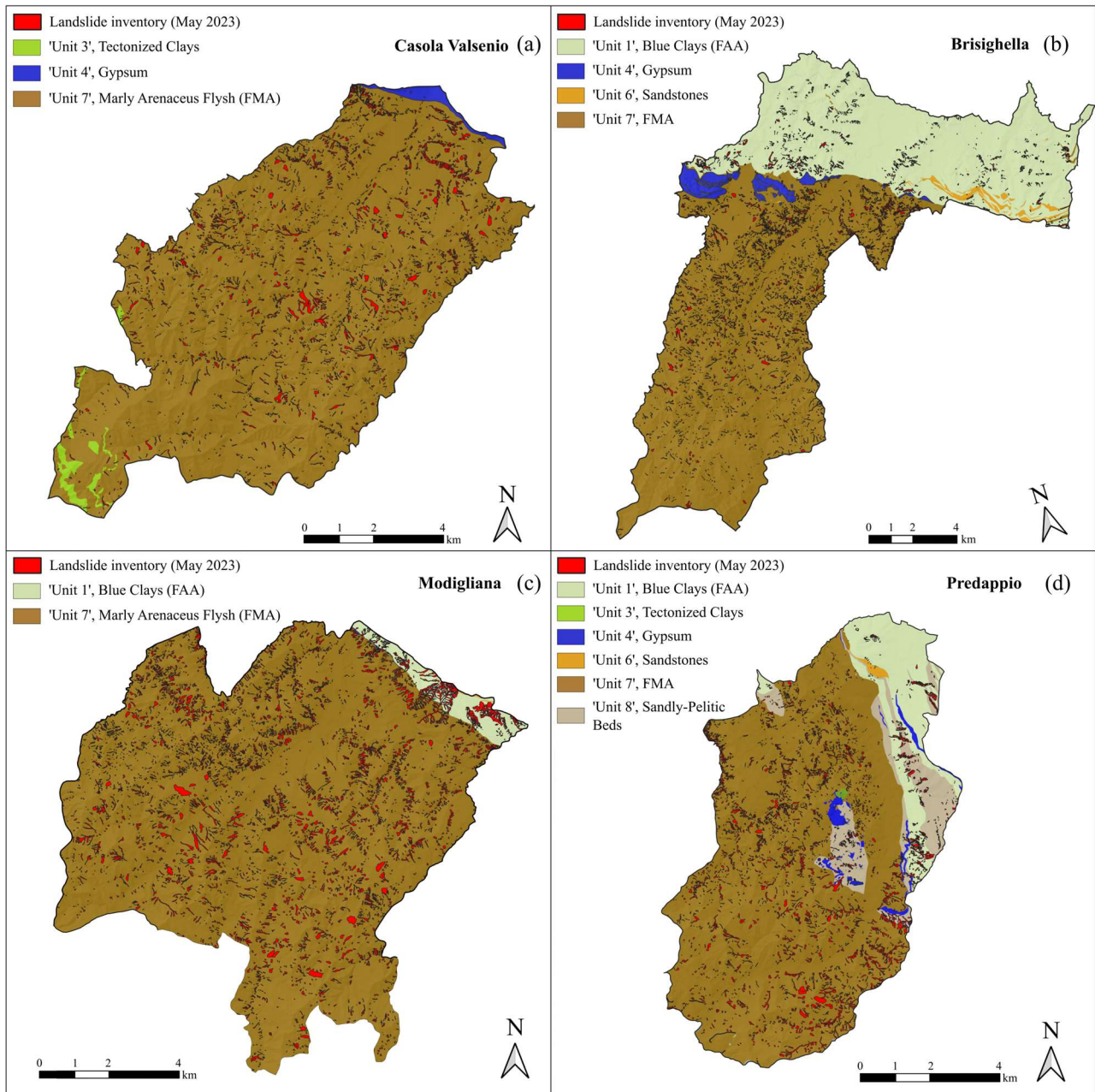
245 Automated methods for landslide mapping were applied in the four municipalities shown in Figure 2
 246 (Casola Valsenio, Brisighella, Modigliana and Predappio). These areas, situated in the eastern part of
 247 the Emilia-Romagna region (Fig. 1), were significantly affected by the meteorological event of May
 248 2023. The Casola Valsenio area was chosen for training models, whereas Brisighella, Modigliana,
 249 and Predappio served as validation sites. In their study, Berti et al. (2026) used Casola Valsenio to
 250 train a conventional U-Net model, and compared the results with those obtained using the traditional
 251 NDVI change detection method. We opted to continue using this area for training to ensure
 252 consistency with previous research. The other three areas were selected because together they
 253 encompass approximately 35% of the regions most impacted by landslides and represent around 25%
 254 of the recorded 80,997 landslide events (20148). Moreover, these municipalities exhibit varied

geological conditions, with both similarities and differences to Casola Valsenio, thereby offering a comprehensive test of the models' adaptability to diverse terrains and landforms.

To represent geological variability in a way that is both meaningful and operational for the analysis, we adopted the lithological classification introduced by Berti et al. (2025), which groups the over 600 geological formations of the Emilia-Romagna region into eight lithological units (Table 2, Fig. 1b). This classification combines lithological characteristics with structural domains, acknowledging that the same rock type may exhibit different mechanical properties depending on its tectonic setting. Units 1 to 3 consist of fine-grained materials (e.g., clays, marls, tectonized shales), while units 4 to 8 mainly comprise coarse-grained rocks (e.g., sandstones, flysch, conglomerates). This simplified scheme captures key differences in the weathering behavior and mechanical response of the rock masses, which are known to influence the type and frequency of landslides. Fine-grained units generally produce cohesive soils prone to earth slides and flows, whereas coarse-grained units generate granular soils more susceptible to debris slides and debris flows. This categorization aligns with the "earth" and "debris" material types defined in the Cruden and Varnes (1996) landslide classification and is used throughout this study to interpret landslide patterns in relation to geological setting.

Table 2. Simplified classification of the geological formations in the Emilia-Romagna region into eight litho-structural units, adapted from Berti et al. (2025). The classification is based on lithological composition and structural domain. Units 1 to 3 predominantly include fine-grained rocks, while units 4 to 8 are composed mainly of coarse-grained lithologies.

Unit ID	Lithology	Domain	Structural position	Geological Age
1	Clays, silty clays, and marly clays	Padano-Adriatic	Outer Foredeep	Pliocene to Pleistocene
2	Marls and marly clays	Epiligurian	Wedge-top basins	Oligocene to Miocene
3	Clay shales, clay breccias, tectonized clays, olistostromes	Ligurian	Accretionary wedge	Cretaceous to Eocene
4	Massive rocks: basalts, serpentines, limestones, arenites	Ligurian, Epiligurian	Accretionary wedge Wedge-top basins	Cretaceous to Miocene
5	Flysch rocks made of rhythmic alternations of sandstones, limestones, pelites, and shales	Ligurian, Epiligurian	Accretionary wedge Wedge-top basins	Cretaceous to Eocene
6	Weakly cemented sandstones and conglomerates	Padano-Adriatic	Outer Foredeep	Pliocene to Pleistocene
7	Flysch rocks made of rhythmic alternations of sandstones and pelites	Tuscan-Umbrian	Inner Foredeep	Miocene
8	Weakly cemented sandstones with interbedded pelitic layers	Padano-Adriatic	Outer Foredeep	Pliocene to Pleistocene



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Figure 2. Schematic geological map of the municipalities of Casola Valsenio (a), Brisighella (b), Modigliana (c) and Predappio (d). FMA = Marnoso-Arenacea Formation, FAA = Blue Clays Formation.

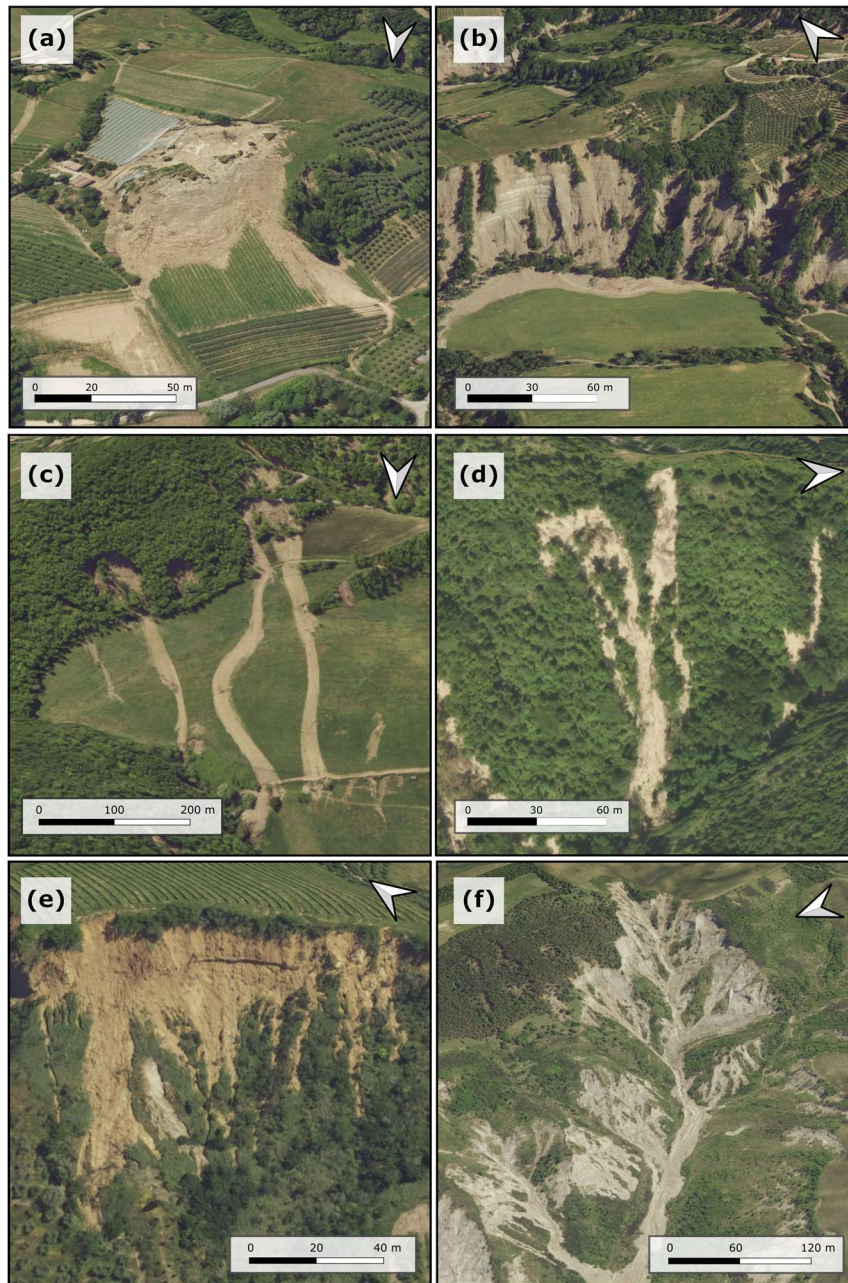


Figure 3. Examples of Landslides in Romagna from May 2023 event following Berti et al. 2025 classification.

- a)** Rockslide (RS) at Baccagnano, Brisighella, triggered during the night of May 2-3. (WGS84: 44.2069085, 11.7739792).
- b)** Debris slide (DS1) at Monte Albano, Casola Valsenio. (WGS84: 44.2356093, 11.6682287).
- c)** Debris flow with extended runout on gentle unforested slope (DF1) at Casola Valsenio. (WGS84: 44.20834, 11.64312).
- d)** Debris flow with limited runout on steep forested slope (DF2) at Modigliana. (WGS84: 44.15124, 11.75511).
- e)** Earth Slide (ES) at Brisighella. (WGS84: 44.23912, 11.78393).
- f)** Earth Flow (EF) at Predappio. (WGS84: 44.08947, 11.99627).

Images © Regione Emilia-Romagna, Geoportale 3D (<https://mappe.regione.emilia-romagna.it>)

4. Methods

4.1 Data

Deep-learning models for landslide mapping typically utilize a multi-layer input approach, integrating a combination of satellite or aerial imagery, topographical data, and land use information to capture the complex characteristics of landslides (Ghorbanzadeh et al., 2019, Meena et al., 2022). In our analysis, we considered the following seven layers (Fig. 4):

- L1) Post-event Sentinel-2 images: Four-band (RGB+NIR) satellite images at 10 m spatial resolution, acquired after the second rainfall event on May 23, 2023. Sentinel-2 Level-2A (L2A) products (bottom-of-atmosphere surface reflectance) were used. They represent the first post-event Sentinel-2 images obtained with minimal cloud cover (Fig. 4b).
- L2) Pre-event Sentinel-2 images: Four-band (RGB+NIR) satellite images at 10 m spatial resolution acquired in May 2022, one year prior to the event. Sentinel-2 Level-2A (L2A) products (bottom-of-atmosphere surface reflectance) were used. These images provide a similar vegetation state to that of May 2023 and were selected due to the lack of cloud-free images in the two months preceding the event (Fig. 4a).
- L3) NDVI change Sentinel-2 map: Single-band raster generated by subtracting the pre-event NDVI map (Kriegler et al., 1969) derived from L2 from the post-event NDVI map created using L1 (Fig. 4c).
- L4) Post-event CGR images: Four bands (RGB+NIR) aerial photographs with a very high resolution of 0.2 m captured on May 23, 2023, shortly after the event. These images were committed by the Emilia-Romagna Region specifically for the disaster (Fig. 4f).
- L5) Pre-event AGEA images: Four bands (RGB+NIR) aerial photographs with a very high resolution of 0.2 m captured 3 years before the event (April to June 2020). These images were committed by the Agency for Agricultural Payments mainly for agricultural purposes (Fig. 4e).
- L6) NDVI change CGR map: Single-band raster generated by subtracting the pre-event NDVI map derived from L5 from the post-event NDVI map created using L4 (Fig. 4g).
- L7) Slope map: Single-band raster derived from the 5x5 m Digital Terrain Model (DTM) of the Emilia Romagna Region. The DTM, which is based on the 1:5000 scale Regional Technical Map, captures the pre-event slope morphology (Fig. 4d).

These layers were combined to define seven data-availability scenarios (cases 1-7 in Table 3), each reflecting realistic operational conditions following a landslide-triggering event.

In all scenarios, the availability of a slope map (L7) was assumed. In the present study, L7 was derived exclusively from the regional 5×5 m DTM. However, this assumption reflects a more general operational condition, as slope information can be readily derived from freely available global digital elevation models (e.g., SRTM, ASTER, ALOS, Copernicus DEM), which are commonly accessible even in data-scarce emergency contexts.

Case 1 (Table 3) represents the most data-limited scenario, where only post-event Sentinel-2 imagery (L1) is available. Owing to its global coverage, short revisit time, and free accessibility, Sentinel-2 data often constitute the primary information source immediately after large-scale disasters (Wasowski et al., 2014; Yang et al., 2019; Ban et al., 2020).

336 Case 2 considers situations in which pre-event Sentinel-2 imagery (L2) is also accessible, allowing
337 the computation of an NDVI change map (L3). While this condition is frequently met, prolonged
338 cloud cover or strong seasonal vegetation variability may prevent the availability of suitable pre-event
339 images.

340 Cases 3 and 4 describe scenarios in which high-resolution post-event imagery is acquired, either alone
341 (case 3) or in combination with Sentinel-2 data (case 4). Such data may originate from very-high-
342 resolution satellite systems (e.g., WorldView, Pleiades, SkySat) or dedicated aerial surveys and
343 substantially improve the detection of small or geomorphically subtle landslides.

344 Cases 5 and 6 address less common situations in which high-resolution imagery is available both
345 before and after the event, optionally complemented by a high-resolution NDVI change map. The
346 effectiveness of these scenarios depends strongly on the temporal proximity of the pre-event imagery
347 to the triggering event.

348 Finally, case 7 represents the most data-rich configuration, combining all available layers (L1-L7).

349 Although the post-event (L4), pre-event (L5), and derived NDVI change (L6) aerial datasets were
350 originally available at 0.2 m spatial resolution, their use at native resolution resulted in systematic
351 out-of-memory errors during training and inference due to the large size of the input tensors. To
352 ensure computational stability, were therefore downsampled to 2 m, representing a practical
353 compromise between spatial detail and computational feasibility for regional-scale mapping. By
354 contrast, Sentinel-2 layers (L1–L3) and the slope layer (L7) were resampled to the same 2 m grid
355 solely to ensure pixel-wise alignment when used together with the aerial datasets; this step does not
356 add information beyond their native resolution and was performed for operational consistency.

357 **Table 3.** Overview of the various input layer configurations used during the training processes. 'NDVI'
 358 refers to the 'ΔNDVI-CGR' Change map created using CGR imagery, while 'ΔNDVI-S2' refers to
 359 the NDVI Change map created using Sentinel-2 imagery.

Cases	Name	Model	Sentinel2 post [L1]	Sentinel2 pre [L2]	Sentinel2 ΔNDVI [L3]	CGR post [L4]	AGEA pre [L5]	CGR ΔNDVI [L6]	Slope [L7]
Case 1	U1	U-Net	✓	X	X	X	X	X	✓
	S1	SegForm	✓	X	X	X	X	X	✓
Case 2	U2	U-Net	✓	✓	✓	X	X	X	✓
	S2	SegForm	✓	✓	✓	X	X	X	✓
Case 3	U3	U-Net	X	X	X	✓	X	X	✓
	S3	SegForm	X	X	X	✓	X	X	✓
Case 4	U4	U-Net	✓	✓	✓	✓	X	X	✓
	S4	SegForm	✓	✓	✓	✓	X	X	✓
Case 5	U5	U-Net	X	X	X	✓	✓	X	✓
	S5	SegForm	X	X	X	✓	✓	X	✓
Case 6	U6	U-Net	X	X	X	✓	✓	✓	✓
	S6	SegForm	X	X	X	✓	✓	✓	✓
Case 7	U7	U-Net	✓	✓	✓	✓	✓	✓	✓
	S7	SegForm	✓	✓	✓	✓	✓	✓	✓

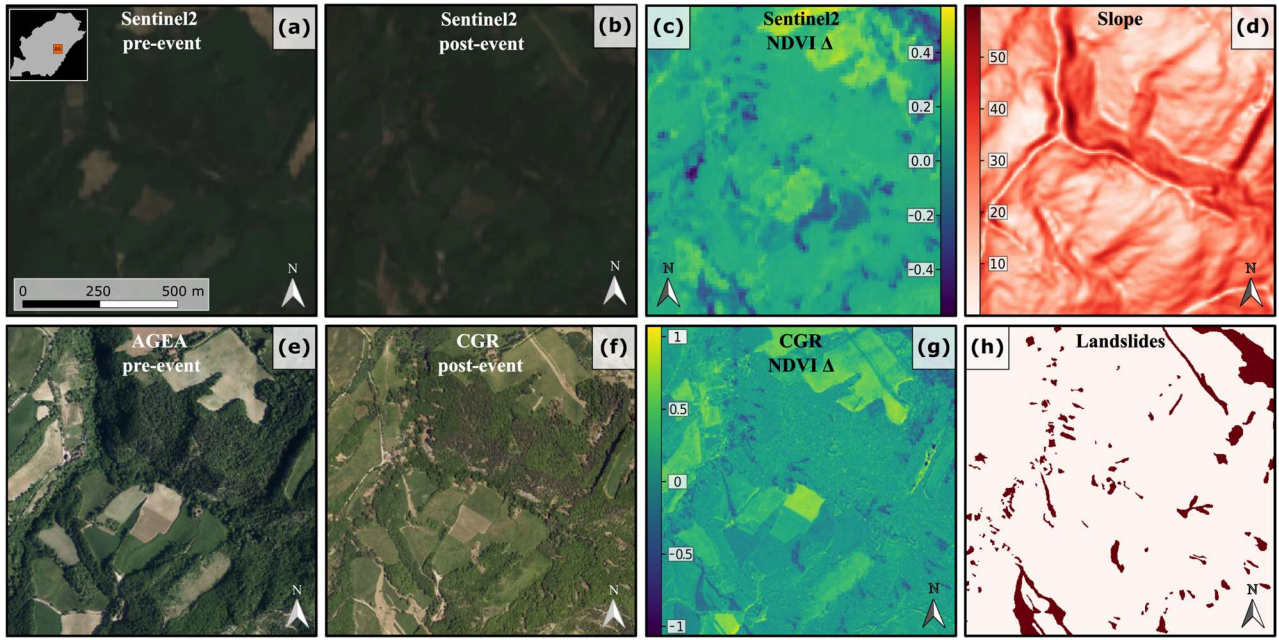


Figure 4. Example of the layers included in case 7, Tile No. 86. (a) Pre-event Sentinel-2 image (2022); (b) Post-event Sentinel-2 image; (c) NDVI change map derived from Sentinel-2 imagery; (d) Slope map; (e) AGEA imagery (2020), pre-event; (f) CGR aerial image (2023), post-event; (g) NDVI change map derived from high resolution imagery; (h) Binary raster showing manually mapped landslides (ground truth).

4.4 Data preparation and training-testing design

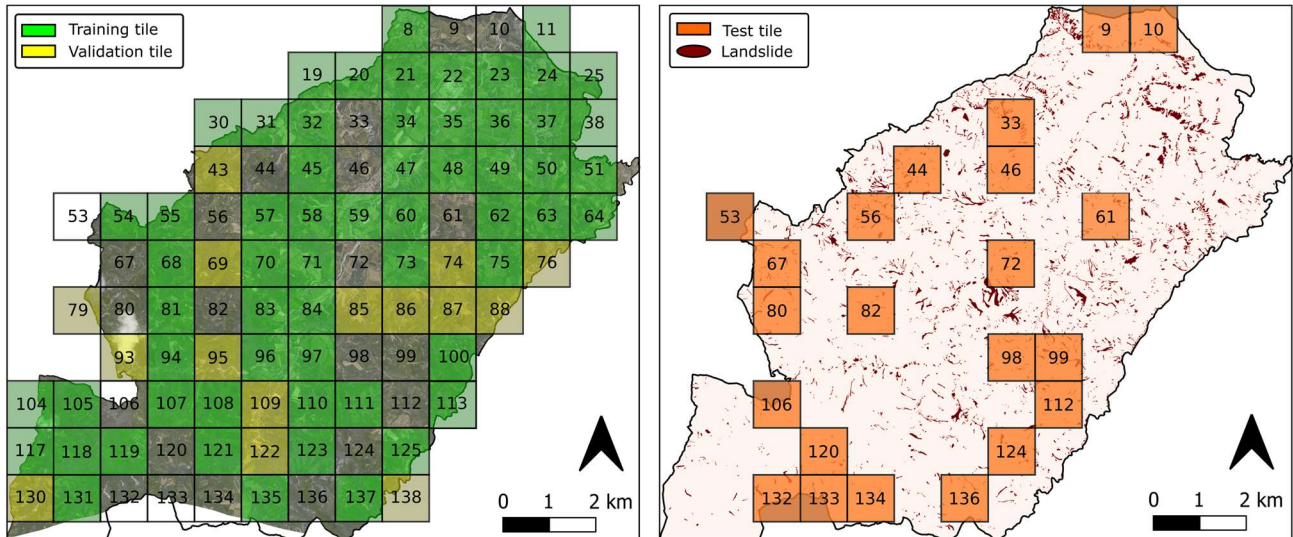
To reproduce operational conditions typical of post-event emergency mapping, the deep-learning workflow was designed to prioritise spatial generalisation rather than random pixel-level splitting. All input layers described in Section 3.1 were harmonised to a common spatial resolution and grid prior to training.

The training area (Casola Valsenio municipality) was subdivided into regular 1×1 km tiles (Fig. 5) to define consistent spatial sampling units and to reduce spatial autocorrelation effects. When resampled to a spatial resolution of 2 m, each tile corresponds to 512×512 pixels in the CGR imagery. Individual tiles typically include a large number of mapped landslides (often exceeding 50), providing a robust representation of landslide morphology and surrounding land cover.

Instead of conventional random splits (e.g. 70/30 or 80/20; Hastie et al., 2009), we adopted a strategy tailored to emergency response scenarios. The models were trained exclusively on data from Casola Valsenio and subsequently applied, without retraining, to three neighbouring municipalities, Predappio, Modigliana, and Brisighella, for independent testing. This setup simulates a realistic operational context in which a model trained on a limited reference area is rapidly deployed to map other affected regions.

Within Casola Valsenio, a stratified random sampling strategy was used to partition the tiles into training, validation, and internal testing subsets. Out of 97 tiles, 60 (62%) were assigned to training, 15 (15%) to validation, and 22 (23%) to internal testing (Fig. 5). External evaluation was then conducted on all tiles from the three remaining municipalities, providing a spatial hold-out test to assess cross-municipality generalisation. This tile-based spatial splitting yields a strongly imbalanced

389 pixel-level segmentation problem, because landslide pixels represent only a small fraction of the
 390 mapped area. In Casola Valsenio (training area), landslide pixels account for $\sim 4\%$ of the total pixels
 391 (non-landslide:landslide ratio $\sim 24:1$). A comparable imbalance characterises the external test areas,
 392 with $\sim 6\%$ landslide pixels in Predappio ($\sim 16:1$), $\sim 7.6\%$ in Modigliana ($\sim 12:1$), and $\sim 5\%$ in
 393 Brisighella FMA ($\sim 19:1$). The most extreme imbalance occurs in Brisighella FAA, where landslide
 394 pixels represent $\sim 2\%$ of the pixels ($\sim 48:1$). The implications of this class imbalance for model
 395 evaluation are discussed in Section 4.7.



396
 397 **Figure 5.** Training and test tiles used for models' construction in Casola Valsenio, illustrating the
 398 spatial split of the area.

399

400 4.5 Deep Learning Semantic Segmentation Models

401 To evaluate the performance of different deep-learning paradigms under variable data-availability
 402 scenarios (Table 3), two semantic segmentation architectures were implemented: a convolutional
 403 encoder-decoder network (U-Net) and a Transformer-based segmentation model (SegFormer).
 404 Schematic representations of both architectures are provided in Fig. 6 to facilitate understanding and
 405 comparison.

406

407 *4.5.1 U-Net*

408 U-Net is a convolutional neural network originally developed for biomedical image segmentation
 409 (Ronneberger et al., 2015) and subsequently adopted in a wide range of geomorphological and Earth-
 410 observation applications. Its encoder-decoder structure enables the extraction of contextual
 411 information while preserving fine-scale spatial detail, a key requirement for accurate landslide
 412 delineation. The suitability of U-Net for landslide mapping has been demonstrated in several recent
 413 studies (Meena et al., 2021, 2022; Ghorbanzadeh et al., 2022, 2023; Nava et al., 2022).

414 The architecture consists of a contracting encoder path and an expanding decoder path. In the encoder,
 415 spatial resolution is progressively reduced through convolutional and max-pooling layers to capture
 416 high-level contextual features. In the decoder, feature maps are upsampled using transposed

417 convolutions and concatenated with corresponding encoder features via skip connections, which
418 preserve spatial information and improve boundary localisation (Fig. 6a).

419 In this study, U-Net was implemented manually in Python and configured for binary semantic
420 segmentation (landslide vs non-landslide). The network employed two downsampling scales, each
421 comprising two convolutional layers. The initial number of filters was set to 64 and increased with
422 network depth to capture progressively more complex features.

423 Training was performed using the Adam optimiser (Kingma and Ba, 2017) with a learning rate of 1
424 $\times 10^{-4}$ and Dice loss, for up to 300 epochs. Early stopping with a patience of 20 epochs was applied
425 based on validation loss. All experiments were conducted using Python 3.8.18, TensorFlow 2.5.0,
426 and CUDA 11.2.

427

428 *4.5.2 SegFormer*

429 SegFormer is a Transformer-based architecture specifically designed for efficient semantic
430 segmentation of high-resolution and multispectral imagery (Xie et al., 2021). It addresses several
431 limitations of earlier Vision Transformer models, including high computational cost, dependence on
432 positional encodings, and inefficient multi-scale feature aggregation (Dosovitskiy et al., 2020).

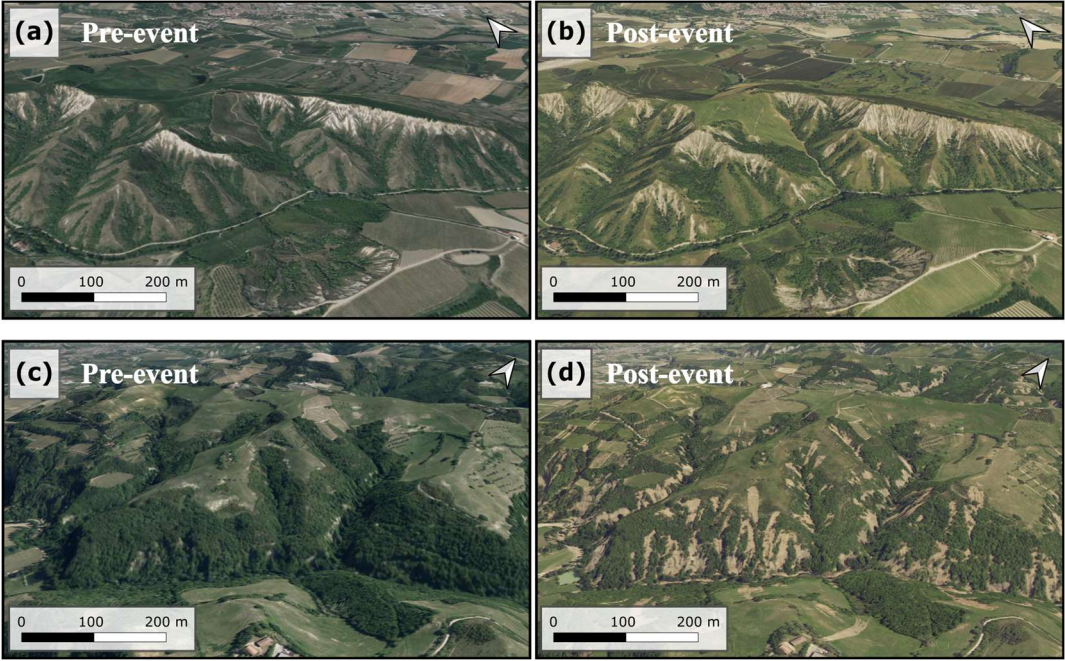
433 The SegFormer architecture comprises two main components (Xie et al., 2021): (i) a hierarchical
434 Transformer encoder that captures features at multiple spatial scales using efficient self-attention
435 without positional encodings and Mix-FFN blocks with depth-wise convolutions, and (ii) a
436 lightweight all-MLP decoder that fuses multi-level encoder features for dense per-pixel prediction
437 (Fig. 6b). The absence of positional encodings allows the model to handle variable input resolutions
438 without interpolation, which is advantageous for heterogeneous landslide imagery.

439 We adopted the MiT-B0 variant (Xie et al., 2021), with hidden sizes [32, 64, 160, 256], encoder
440 depths [2, 2, 2, 2], and a decoder hidden size of 256. Although SegFormer was originally designed
441 for three-channel RGB imagery, we exploited the flexibility of the
442 SegformerForSemanticSegmentation implementation provided by the Hugging Face Transformers
443 library (Wolf et al., 2020) to accommodate multi-channel inputs.

444 Training was performed using Cross-Entropy loss and the Adam optimiser with a learning rate of 1
445 $\times 10^{-3}$, for up to 300 epochs, with early stopping based on validation loss (patience = 20 epochs).
446 Computations were carried out using Python 3.8.10, PyTorch 1.9.0 (Paszke et al., 2019), and CUDA
447 11.1.

464 Arenacea Formation (Brisighella FMA) and one characterised by Pliocene Blue Clays (Brisighella
465 FAA; Fig. 3d).

466 The FAA domain is dominated by earth flows (EF) and earth slides (ES), typically developed within
467 fine-grained, clay-rich materials and often associated with badland-like morphologies. These
468 landslides generally exhibit elongated shapes, lobate toes, and sparse or degraded vegetation cover,
469 reflecting repeated reactivation processes (Fig. 7a-b). In contrast, landslides in the FMA domain, both
470 in Brisighella and in the training area, are mainly debris slides (DS) and debris flows (DF) triggered
471 within thin colluvial layers overlying flysch bedrock. These failures tend to be rapid, with well-
472 defined scarps and runout zones, and commonly affect densely vegetated slopes (Fig. 7c-d).



473
474 **Figure 7.** Comparison examples are shown between pre-event (a) and post-event (b) earth flow (EF)
475 landslides in the Blue Clays Formation (FAA) and pre-event (c) and post-event (d) debris slide (DS)
476 in the Marnoso-Arenacea Formation (FMA), both triggered by the May 2023 disaster in Brisighella.
477 Images © Regione Emilia-Romagna, Geoportale 3D (<https://mappe.regione.emilia-romagna.it>)

478
479 By explicitly separating these lithological contexts, we evaluated whether models trained solely on
480 FMA-type landslides could generalise to areas characterised by fundamentally different landslide
481 processes, geometries, and surface expressions. A summary of the spatial extent of each lithological
482 unit within the test municipalities is provided in Table 4 to support interpretation of the evaluation
483 results.

484
485
486
487
488

Table 4. Overview of the 1 km² tiles used for model training, validation, and testing across the analysed municipalities. The table reports the number of tiles assigned to each dataset split and their distribution within the main lithological units, namely the Marnoso-Arenacea Formation (FMA) and the Pliocene Blue Clays (FAA). External testing was performed on 100% of the tiles from Predappio, Modigliana, and Brisighella to assess cross-municipality and cross-lithology model generalisation.

Municipality	Total tiles [1 km ²]	Training	Validation	Internal test	External test	FMA 'Unit 7'	FAA 'Unit 1'
Casola Valsenio	97	60	15	22	–	97	0
Predappio	90	–	–	–	90	75	15
Modigliana	105	–	–	–	105	99	6
Brisighella FMA	116	–	–	–	116	116	0
Brisighella FAA	80	–	–	–	80	0	80

4.7 Evaluation Metrics and Expert Judgment

4.7.1 Quantitative Metrics

To assess the performance of our landslide detection models, we employed two widely used metrics in semantic segmentation tasks: the F1 score and the Intersection over Union (IoU). These metrics were calculated on the entirety of the test images, rather than averaging scores across individual cells, to provide a comprehensive evaluation of the models' performance.

The F1 score (Dice coefficient) is the harmonic mean of precision and recall (Tharwat, 2020). We report F1 alongside IoU because the pixel-wise segmentation task is strongly imbalanced (Section 4.4), with landslide pixels representing only a small fraction of the mapped area. The F1 score is calculated as:

$$F1 = 2 \cdot \frac{Precision \cdot Recall}{Precision + Recall} \quad (1)$$

where Precision = TP / (TP + FP) and Recall = TP / (TP + FN), with TP, FP, and FN representing True Positives, False Positives, and False Negatives, respectively.

The Intersection over Union (IoU), also known as the Jaccard index, is another common metric for assessing segmentation accuracy (Rezatofighi et al., 2019). It measures the overlap between the predicted segmentation mask and the ground truth mask. The IoU is calculated as:

$$IoU = \frac{TP}{TP + FP + FN} \quad (2)$$

Both metrics range from 0 to 1, with 1 indicating perfect prediction. By using these metrics on the entire test images, we can evaluate how well our models perform in detecting and delineating landslides across varied landscapes and geological contexts.

4.7.2 Expert Judgment

In automated landslide mapping, a higher algorithmic score does not necessarily correspond to a higher quality map (Zhang et al., 2018; Isensee et al., 2021). This issue is particularly relevant in automated landslide mapping, where the practical utility of the map often depends more on its interpretability and coherence with geomorphological reality than on its statistical performance alone.

For example, a model might overpredict large landslide areas to optimize pixel-wise metrics like IoU, while neglecting smaller or fragmented features that are critical for field validation and emergency response. Conversely, models may introduce noise or irregular boundaries that artificially increase recall but reduce the practical readability of the map.

While this issue is typically addressed by selecting evaluation metrics that emphasize specific components of the confusion matrix (e.g., recall over precision), in emergency scenarios not all components carry the same operational weight. For instance, in crisis management, a higher recall may be preferred to avoid missing active landslides, even if this comes at the cost of more false positives. In contrast, in densely urbanized areas or during resource-constrained interventions, precision may become more important to avoid unnecessary alarms and misallocation of efforts.

To address this challenge, three of the authors in this study provided their expert judgment on the quality of the automatically generated landslide maps. This evaluation was conducted without prior knowledge of the scores obtained by these maps or the training input layers used to generate them.

The assessment focused on seven common error types frequently encountered in automated landslide mapping:

E1 - False positives in plowed fields, where plowed land is mistakenly classified as landslide;

E2 - False positives along roads and in urban areas, where the absence of vegetation may be misinterpreted as landslide activity;

E3 - False positives along riverbeds, often due to confusion with recent sediment scouring or bedload deposits;

E4 - False negatives in areas with known landslides;

E5 - Over segmentation, referring to the fragmentation of a single landslide into multiple small polygons;

E6 - Inaccurate delineation of landslide boundaries;

E7 - Omission errors in shadowed areas, such as steep, shaded slopes or regions obscured by cloud cover.

For each error type, the experts assigned a score of 0 (limited error), 1 (moderate error), or 2 (relevant error) to indicate the severity of the issue. These scores were then summed and normalized to calculate an Expert-based Performance Index (EPI):

$$EPI = \frac{14 - \sum_{i=1}^{i=7} E_i}{14}$$

The index ranges from 0 to 1, with 1 representing the highest mapping performance. This evaluation was independently carried out by each of the three experts for all 16 cases included in the analysis (see Table 4). The experts assessed the model cases without prior knowledge of the specific configurations, thereby minimizing potential bias.

558 4.8 Identification of At-Risk Buildings

559 In the context of emergency mapping, a key objective was to identify buildings at risk due to the
560 landslides triggered by the May 2023 event. This was a critical task for the Civil Protection as it
561 allowed for timely intervention and damage assessment. Given the extensive spatial distribution of
562 the damage caused by the event, it was essential to develop a systematic method to detect and assess
563 the buildings at risk. Additionally, accurate damage estimation was necessary to request financial
564 support under the European Union Solidarity Fund (EUSF), adhering to the EU's required timelines
565 for funding requests.

566 To evaluate the effectiveness of our automated landslide mapping models in identifying at-risk
567 buildings, we created a 20-meter buffer around the predicted landslide boundaries. This buffer was
568 compared to the list of buildings that the Civil Protection had classified as "at risk," based on manual
569 mapping (Pizziolo et al., 2024; Berti et al., 2025). The buildings identified by the automated models
570 were then compared to these reference buildings, allowing us to assess the models' capacity for
571 accurate identification. The confusion matrix was used to quantify the performance of the models in
572 identifying buildings at risk, including both false positives (incorrectly flagged buildings) and false
573 negatives (missed buildings).

574 This methodology ensured that the models were tested in the real-world context of emergency
575 response, where timely and accurate damage assessments are crucial for effective resource allocation
576 and intervention. The results of this methodology are discussed further in Section 6.3, while the
577 quantitative outcomes of our models are presented in Section 5.3.

578

579 **5. Results**

580 5.1 Model results and performance

581 The results of the analysis are reported in Table 5, which summarizes the F1 and IoU scores obtained
582 by comparing the automated landslide maps generated by the two deep-learning models (U = U-Net;
583 S = SegFormer) with the manually mapped reference inventory. Seven combinations of input layers
584 (cases 1-7; Table 1) were tested, representing progressively richer input information. The models
585 were trained using data from Casola Valsenio and subsequently applied to four independent test areas:
586 Brisighella FMA, Brisighella FAA, Modigliana and Predappio. In total, 56 automated landslide maps
587 were produced (7 input configurations $\times$ 4 test areas $\times$ 2 models).

588 Across all municipalities, the highest performance was obtained by the most information-rich
589 configurations, particularly U7 and U6 for U-Net and S7 for SegFormer. In Casola Valsenio, the best
590 configuration reached F1 = 0.73 and IoU = 0.57 (U7), while in Brisighella FMA, Brisighella FAA,
591 Modigliana and Predappio the highest F1-scores were consistently around 0.60–0.63 (e.g., U7 = 0.56–
592 0.63; S7 = 0.60–0.61). Performance was systematically lower in Brisighella FAA, where the best
593 configurations reached F1 = 0.53 and IoU = 0.36 (U6/U7), and F1 = 0.52 and IoU = 0.35 (S7) (Table
594 5).

595 Despite relying on a reduced set of inputs, the Sentinel-2-based configuration S2 (post-event Sentinel-
596 2 plus Sentinel-2 NDVI change) achieved competitive results across municipalities. For instance, in
597 Casola Valsenio S2 reached F1 = 0.62 and IoU = 0.45, compared with F1 = 0.73 and IoU = 0.57 for
598 the best-performing configuration (U7). In Brisighella FMA, Modigliana and Predappio, S2 yielded

599 F1-scores between 0.54 and 0.57 (IoU = 0.37–0.40), which are close to the values obtained by the
600 strongest configurations (typically within ~0.03–0.06 in F1 and ~0.03–0.06 in IoU). In Brisighella
601 FAA, S2 remained among the most stable configurations (F1 = 0.47; IoU = 0.31), whereas some
602 intermediate cases showed marked drops (e.g., U3: F1 = 0.29; IoU = 0.17).

603 Overall, increasing the number of input layers resulted in measurable but limited performance
604 improvements. The largest gain was associated with the inclusion of NDVI change information (case
605 2 vs. case 1), while subsequent additions of high-resolution inputs generally produced smaller
606 incremental changes. These results indicate that streamlined configurations based on Sentinel-2 data
607 (case 2), optionally complemented by high-resolution imagery, can achieve segmentation
608 performance close to that obtained using the full set of available inputs (Table 5).

609

610 **Table 5.** F1-score and Intersection over Union (IoU) obtained by the two deep-learning semantic
611 segmentation models (U = U-Net; S = SegFormer) for the seven input-layer configurations (cases 1–
612 7; Table 3), evaluated against the manually mapped reference inventory. Models were trained in
613 Casola Valsenio and applied without retraining to the external test municipalities (Brisighella FMA,
614 Brisighella FAA, Modigliana, Predappio). The table therefore reports 56 model outputs (7 cases
615 × 4 municipalities × 2 models).

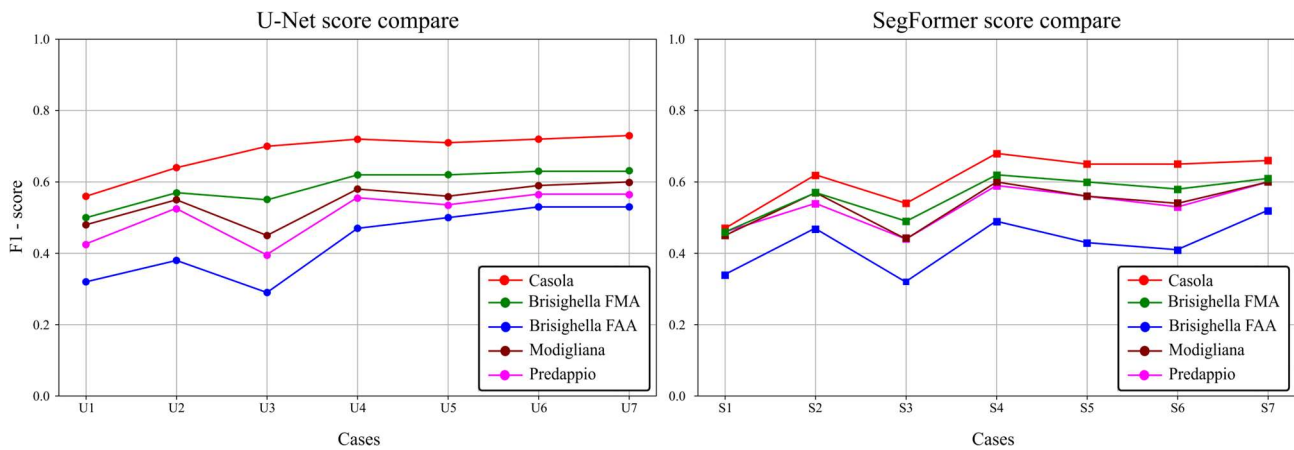
Cases	Models' Name	Casola Valsenio		Brisighella FMA		Brisighella FAA		Modigliana		Predappio	
		F1	IoU	F1	IoU	F1	IoU	F1	IoU	F1	IoU
Case 1	U1	0.56	0.39	0.50	0.33	0.32	0.19	0.48	0.31	0.42	0.27
	S1	0.47	0.31	0.46	0.30	0.34	0.21	0.45	0.29	0.46	0.30
Case 2	U2	0.64	0.47	0.57	0.40	0.38	0.24	0.55	0.38	0.52	0.35
	S2	0.62	0.45	0.57	0.40	0.47	0.31	0.57	0.40	0.54	0.37
Case 3	U3	0.70	0.54	0.55	0.38	0.29	0.17	0.45	0.29	0.39	0.24
	S3	0.54	0.37	0.49	0.32	0.32	0.19	0.44	0.29	0.44	0.28
Case 4	U4	0.72	0.56	0.62	0.45	0.47	0.31	0.58	0.41	0.55	0.38
	S4	0.68	0.52	0.62	0.45	0.49	0.33	0.60	0.43	0.59	0.42
Case 5	U5	0.71	0.55	0.62	0.45	0.50	0.34	0.56	0.39	0.53	0.36
	S5	0.65	0.48	0.60	0.43	0.43	0.28	0.56	0.39	0.56	0.39
Case 6	U6	0.72	0.56	0.63	0.46	0.53	0.36	0.59	0.42	0.56	0.38
	S6	0.65	0.48	0.58	0.41	0.41	0.26	0.54	0.37	0.53	0.36
Case 7	U7	0.73	0.57	0.63	0.46	0.53	0.36	0.60	0.42	0.56	0.39
	S7	0.66	0.49	0.61	0.44	0.52	0.35	0.60	0.43	0.60	0.43

616

Figure 8 illustrates the evolution of F1-scores across the seven input configurations for each municipality. For both models, the inclusion of NDVI change information (case 2) led to an increase in F1-score relative to case 1. In Casola Valsenio, F1 increased from 0.56 to 0.70 for U-Net and from 0.47 to 0.54 for SegFormer. In Modigliana and Brisighella FAA, the corresponding increases were approximately +0.15 to +0.16 for U-Net. Beyond case 2, U-Net showed gradual increases in F1-score of 0.01-0.04 per configuration up to case 7, whereas SegFormer exhibited smaller variations, generally within ± 0.02 , indicating a lower sensitivity to further input enrichment.

Model performance varied across municipalities. In Casola Valsenio, U-Net F1-scores increased from 0.56 (U1) to 0.73 (U7), while SegFormer ranged from 0.47 (S1) to 0.66 (S7). In Predappio, Modigliana, and Brisighella FMA, U-Net F1-scores were between 0.55 and 0.63, whereas SegFormer values remained more stable, typically between 0.55 and 0.60. Brisighella FAA consistently showed the lowest performance for both models: U-Net F1-scores ranged from 0.29 (U3) to 0.53 (U7), and SegFormer values ranged from 0.32 to 0.52, with no systematic improvement beyond case 2. IoU values exhibited the same spatial pattern, decreasing from 0.35-0.46 in FMA-dominated areas to 0.17-0.36 in the FAA domain.

632



633

Figure 8. F1-score comparison across the seven input-layer configurations for U-Net (left) and SegFormer (right). The numbering matches the input-layer cases defined in Table 3 (e.g., U1 and S1 refer to Case 1; U2 and S2 refer to Case 2), and values are summarised in Table 5.

Figures 9 and 10 complement the quantitative results by providing a visual interpretation of model performance under contrasting conditions. Figure 9 illustrates representative cases in which both models performed well, showing a strong spatial agreement between the manually mapped inventory and the automated outputs. In these panels, the overlap between the reference landslide map (blue) and the predicted map (yellow) produces a beige colour, corresponding to true positives and indicating accurate landslide delineation. These examples mainly refer to debris slides and debris flows occurring in well-illuminated areas with widespread vegetation removal, where spectral and textural changes are pronounced.

Figure 10, by contrast, focuses on more challenging scenarios, where the models exhibited recurrent error patterns. In Brisighella FMA, false positives (yellow polygons) are frequently mapped along riverbanks and on recently constructed buildings, likely associated with flood-related sediment

648 redistribution and strong spectral contrasts. The most critical conditions are observed in Brisighella
649 FAA, where widespread false positives occur within badland morphologies developed on Blue Clay
650 formations, highlighting the limited transferability of models trained predominantly on flysch-
651 dominated terrains. In Modigliana, false positives (yellow polygons) are often observed in plowed
652 fields, where agricultural disturbance produces spectral signatures similar to recent landslides.
653 Finally, in Predappio, false negatives (blue polygons) dominate in shadowed slopes, where reduced
654 illumination and altered colour profiles limit the detectability of landslide features.

655 By explicitly contrasting successful detections (Figure 9) with failure modes (Figure 10), these visual
656 examples clarify how different surface conditions, illumination effects, and lithological settings
657 influence model performance, supporting the quantitative trends reported in Table 5 and Figure 8.

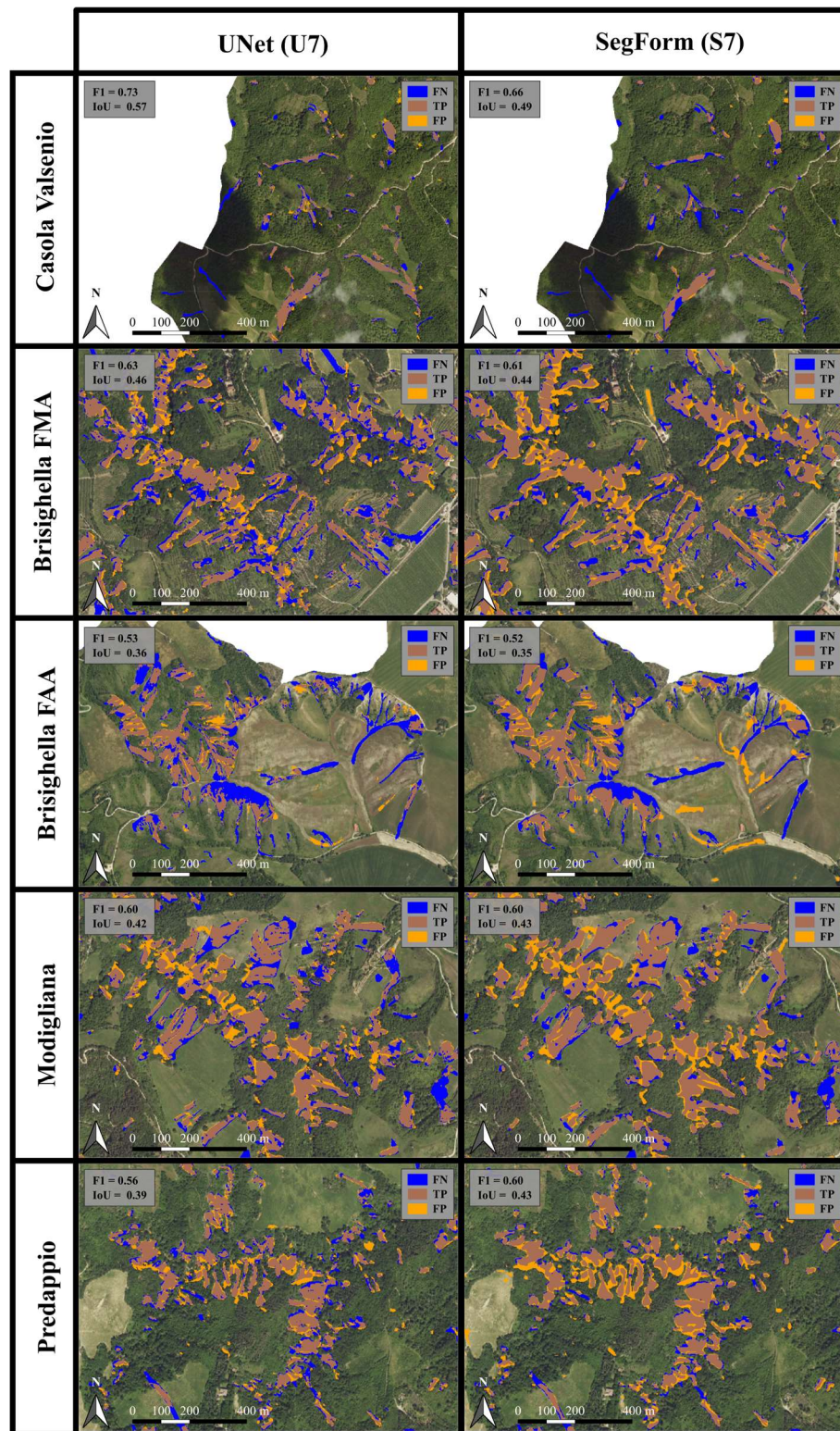


Figure 9. Comparative analysis of landslide mapping results using the U7 (U-Net) and S7 (SegFormer) configurations across the analysed municipalities (Casola Valsenio, Brisighella FMA, Brisighella FAA, Modigliana and Predappio). The figure shows the spatial distribution and extent of the mapped landslides for both models. In the overlays, the reference inventory (True map) is shown in blue and the model prediction in yellow; their intersection is shown in brown (true positives), while blue-only and yellow-only areas correspond to false negatives and false positives, respectively.

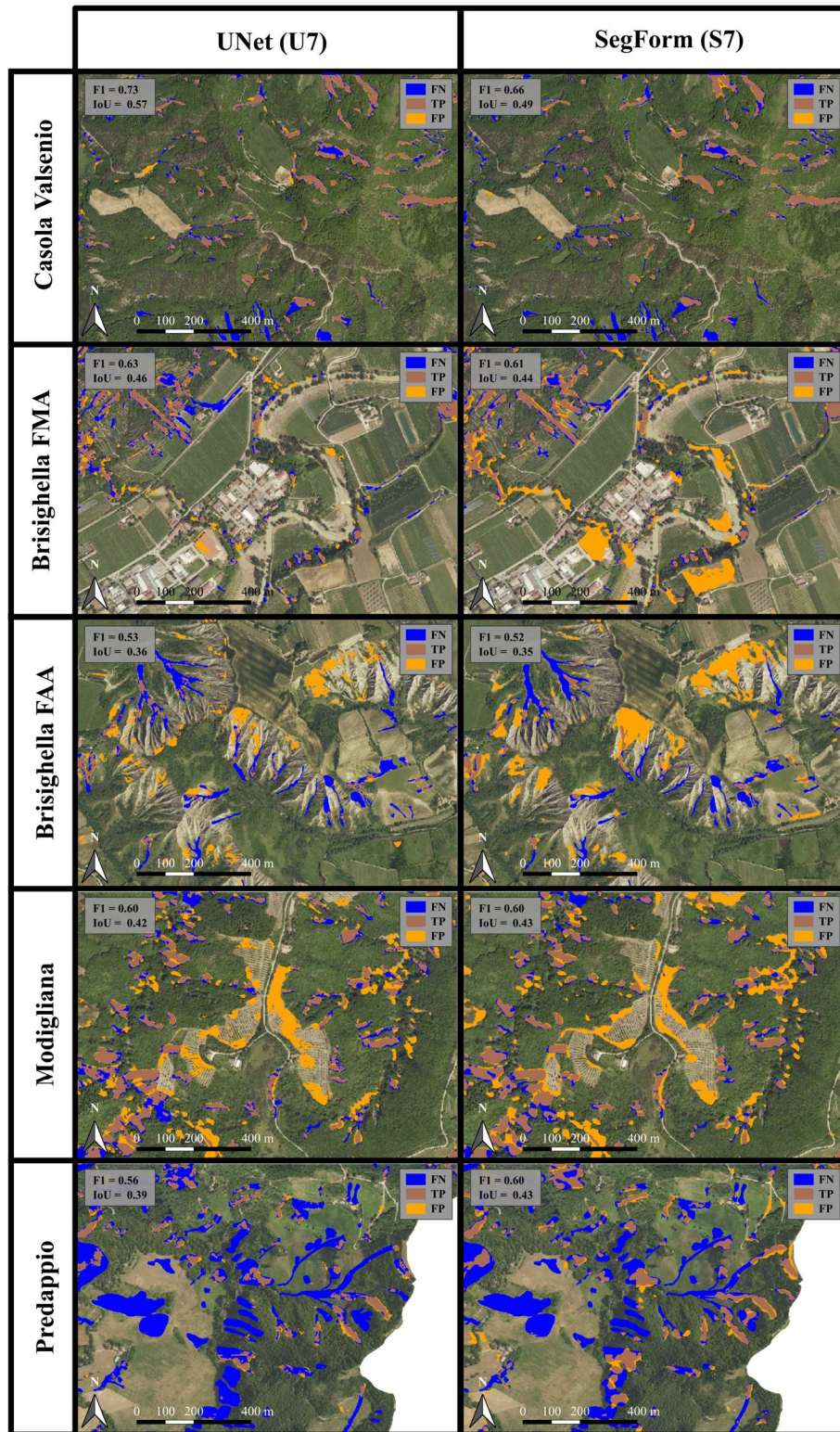


Figure 10. Representative examples of challenging scenarios for landslide detection using the U7 (U-Net) and S7 (SegFormer) configurations across municipalities. The panels highlight recurrent error patterns under various environmental conditions: Brisighella FMA shows riverbanks and buildings as False Positives (FP), Brisighella FAA displays both False Positives and False Negatives (FP & FN) in badlands, Modigliana FN is observed on plowed fields, and Predappio FN occurs under shadows.

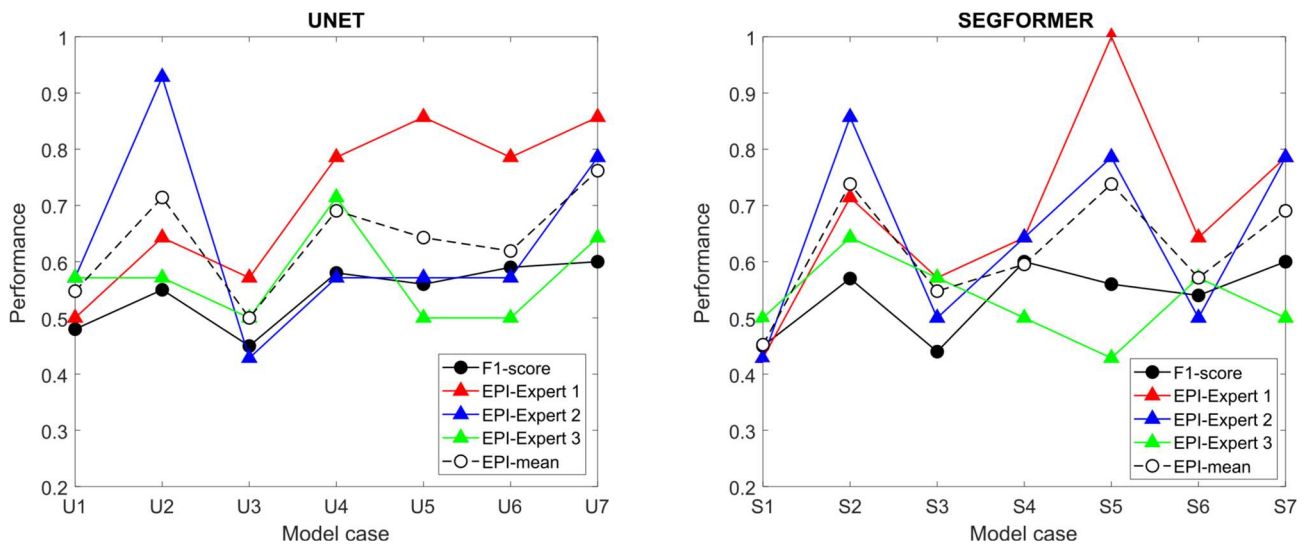
5.2 Expert judgement

Figure 11 compares the quantitative performance (F1-score) of all model configurations with the Expert Performance Index (EPI) derived from the independent assessment of three experts (Section 3.7). The EPI values reflect the severity scores assigned to seven recurrent mapping errors (E1–E7), which were aggregated and normalized to obtain a single index per model configuration.

Across configurations, EPI values show the same overall ranking observed for the F1-scores (Fig. 11). While the F1-scores span a relatively narrow interval across cases (Table 5), the EPI values exhibit a comparable monotonic increase from simpler to more information-rich input configurations, indicating that the expert-based assessment captures the same performance gradient. In particular, configuration U3 is associated with one of the lowest EPI values and one of the lowest F1-scores, whereas U7 consistently falls within the top-performing group in both metrics (Fig. 11).

Differences among individual experts are evident for some intermediate configurations. Specifically, the highest EPI score for Expert 1 is assigned to S5, for Expert 2 to U2, and for Expert 3 to U4, despite these configurations not being uniformly ranked at the top by the other experts or by the F1-scores. However, when expert scores are averaged (mean EPI; dashed line in Fig. 11), the resulting trend closely matches the F1-score curve, reducing the influence of individual preferences and providing a more robust overall assessment of map quality.

690



691

Figure 11. Comparison of model performance based on F1-scores and expert-based evaluations from three independent experts.

694

5.3 Identification of At-Risk Buildings

The identification of at-risk buildings is a crucial component in emergency response scenarios. In the aftermath of the May 2023 disaster, buildings located within landslide boundaries or within a 20-meter buffer zone were considered at risk, influencing subsequent evaluations carried out by the relevant authorities. To assess the effectiveness of our automated landslide mapping for this purpose, we compared the buildings identified by our models to those derived from manual mapping (Pizziolo et al., 2024; Berti et al., 2025).

701

702 Figure 12 presents the confusion matrix for the 14 model configurations, showing the comparison
703 between the buildings flagged by our models and those mapped manually. The F1-scores obtained
704 for each case highlight the models' ability to identify structures at risk. For Case 1, U-Net produced
705 an F1-score of 0.53, while SegFormer performed slightly better at 0.67. The highest F1-scores were
706 achieved in Case 7, with U-Net reaching 0.79 and SegFormer 0.75. Notably, models U2, U4, S6, and
707 U7 consistently yielded the best results, with F1-scores averaging around 0.78, demonstrating their
708 effectiveness in accurately identifying buildings at risk.

709 Despite the relatively high F1-scores, the models still produced both false positives (incorrectly
710 flagged buildings) and false negatives (missed buildings), which underscores the necessity of manual
711 verification in scenarios where precision is critical. However, the performance of the automated
712 models is considerably higher than the manual mapping, especially when considering the large-scale,
713 time-sensitive nature of emergency response. These results suggest that the automated models can
714 significantly assist in the rapid identification of at-risk structures, potentially saving valuable time in
715 the early stages of disaster response. These results are discussed in Section 6.3.

Input Layers		U-NET		SegForm			
Case 1	- Post-event S2 - Slope	U1	F1: 0.53 Predicted value		S1	F1: 0.67	
		Outside	TN 1822 FP 781	Actual value	Outside	TN 2221 FP 382	
Inside	FN 139 TP 515	Inside	FN 137 TP 517				
		Outside		Outside		Inside	
Case 2	- Post-event S2 - Pre-event S2 - Δ NDVI-S2 - Slope	U2	F1: 0.78		S2	F1: 0.77	
		Outside	TN 2455 FP 148	Outside	TN 2421 FP 182		
Inside	FN 144 TP 510	Inside	FN 134 TP 520				
		Outside		Outside		Inside	
Case 3	- Post-event CGR - Slope	U3	F1: 0.72		S3	F1: 0.68	
		Outside	TN 2497 FP 106	Outside	TN 2261 FP 342		
Inside	FN 230 TP 424	Inside	FN 137 TP 517				
		Outside		Outside		Inside	
Case 4	- Post-event S2 - Pre-event S2 - Δ NDVI S2 - Post-event CGR - Slope	U4	F1: 0.78		S4	F1: 0.73	
		Outside	TN 2415 FP 188	Outside	TN 2285 FP 318		
Inside	FN 116 TP 538	Inside	FN 90 TP 564				
		Outside		Outside		Inside	
Case 5	- Post-event CGR - Pre-event AGEA - Slope	U5	F1: 0.76		S5	F1: 0.75	
		Outside	TN 2387 FP 216	Outside	TN 2377 FP 226		
Inside	FN 117 TP 537	Inside	FN 123 TP 531				
		Outside		Outside		Inside	
Case 6	- Post-event CGR - Pre-event AGEA - Δ NDVI CGR - Slope	U6	F1: 0.73		S6	F1: 0.78	
		Outside	TN 2290 FP 313	Outside	TN 2436 FP 167		
Inside	FN 97 TP 557	Inside	FN 125 TP 529				
		Outside		Outside		Inside	
Case 7	- Post-event S2 - Pre-event S2 - Δ NDVI-S2 - Post-event CGR - Pre-event AGEA - Δ NDVI-CGR - Slope	U7	F1: 0.79		S7	F1: 0.75	
		Outside	TP 2436 FP 167	Outside	TP 2282 FP 321		
Inside	FN 120 TN 534	Inside	FN 71 TN 583				
		Outside		Outside		Inside	

Figure 12. Comparison of buildings identified as at risk (within landslide boundaries or within a 20-meter buffer) by automated mapping methods and manual mapping (ground truth), showing the confusion matrices for all five cases evaluated.

720 6. Discussion

721 6.1 Models comparison

722 One of the key observations from this study concerns the differing sensitivity of U-Net and
723 SegFormer to input data combinations, particularly in the context of emergency response mapping.
724 U-Net's performance varies significantly depending on the layers used, while SegFormer yields more
725 consistent results across different configurations. This contrast is likely rooted in their architectural
726 differences. U-Net, designed for pixel-wise segmentation, relies heavily on the spatial and spectral
727 characteristics of the input data, making it more sensitive to the quality and availability of
728 supplementary layers such as pre-event imagery or NDVI change maps. SegFormer, being a
729 transformer-based model, is better equipped to capture global context, thus reducing dependency on
730 specific input combinations and improving robustness under complex geological conditions.

731 This architectural contrast becomes even more evident when models are transferred across different
732 geological domains. As shown in the results (Figure 8), U-Net exhibits a marked decline in
733 performance when moving from areas geologically similar to the training site (e.g., Predappio,
734 Modigliana) to Brisighella FAA, where lithological conditions differ significantly. This suggests that
735 U-Net has limited generalization capacity beyond the Marnoso-Arenacea Formation (FMA), and
736 struggles in fine-grained, clay-rich terrains like those found in the FAA domain. SegFormer, although
737 generally more stable, also underperforms in Brisighella FAA, highlighting the broader challenge of
738 transferring models across heterogeneous geological settings.

739 The feature importance values for U-Net and SegFormer (Figure 13) reveal notable differences in
740 how each model prioritizes the input channels. Upon examining the results, we can observe the
741 following key points:

- 742 ▪ SegFormer places the highest importance on the post-event CGR channels, particularly the
743 Red and NIR bands, followed by Sentinel-2 Red and Sentinel-2 NIR. This makes sense in the
744 context of post-event analysis, as Red and NIR bands are sensitive to vegetation and surface
745 changes, which are crucial for detecting variations in the landscape after a landslide event. The
746 reliance on these bands is likely due to their ability to capture the shift in vegetation cover
747 (from green to beige or brown) after a disturbance, which can be a key indicator of landslide
748 activity. The CGR (post) channels are particularly effective at highlighting these changes,
749 making them essential for SegFormer's decision-making process.
- 750 ▪ U-Net, on the other hand, follows a similar pattern but with lower variability in feature
751 importance across different channels. While the CGR (post) channels remain important, U-Net
752 seems to show a more uniform distribution of feature importance across the input channels,
753 indicating that the model is less sensitive to specific features like Red and NIR. This could be
754 a result of the model's pixel-wise segmentation approach, which might focus more on spatial
755 relationships within the image rather than relying heavily on spectral differences between the
756 bands.

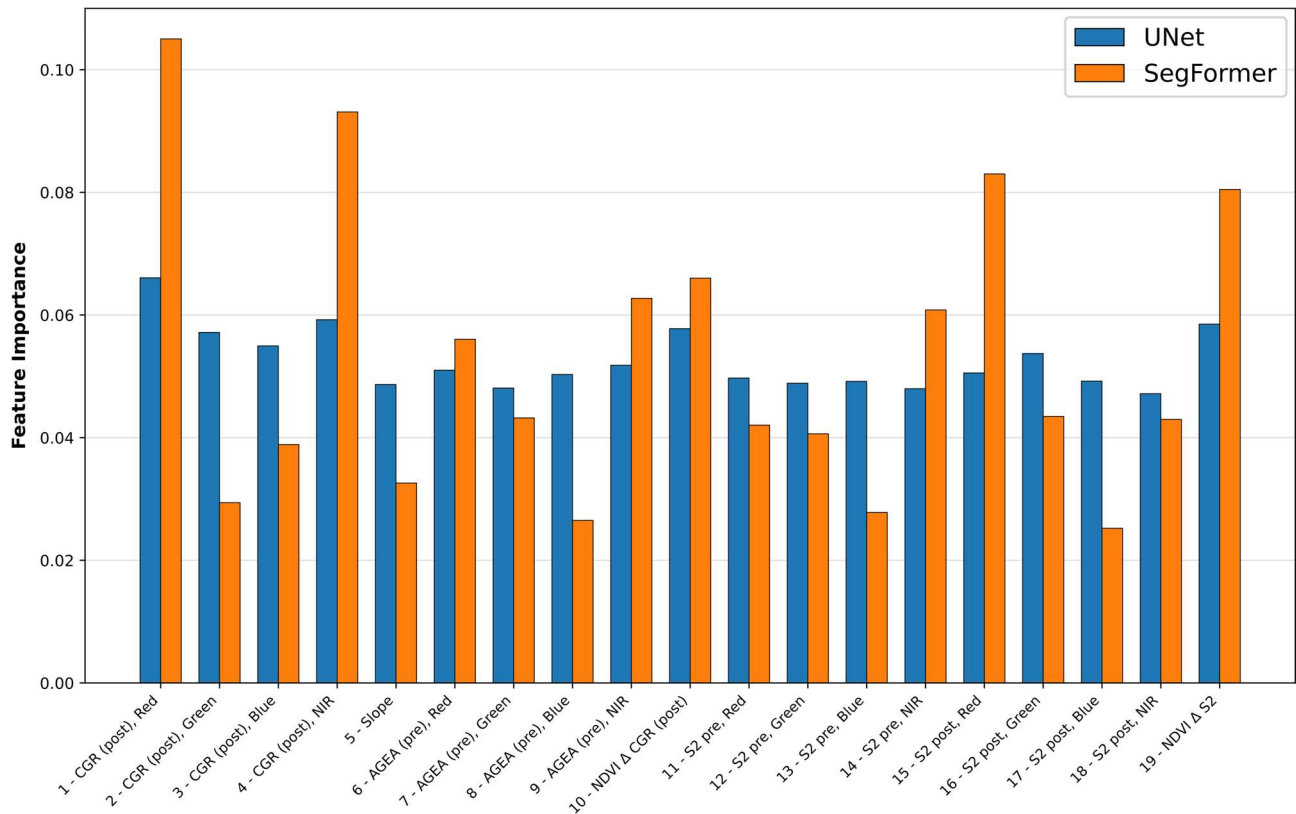


Figure 13. Feature importance values for U-Net and SegFormer across different input channels. The bar chart compares the relative importance assigned to each channel by both models. The blue bars represent U-Net, while the orange bars represent SegFormer. The analysis shows that SegFormer tends to assign higher importance to the post-event CGR channels, particularly the Red and NIR bands, while U-Net exhibits more uniform importance across the channels, with a slightly more localized focus. This difference in feature importance highlights the varying strengths of the models in utilizing spectral information.

Both models exhibit a similar trend in emphasizing the post-event data, particularly the CGR (post) channels. However, SegFormer seems to make use of a wider range of channels, including Sentinel-2 Red and Sentinel-2 NIR, which suggests that SegFormer is more adept at capturing a broader, global context. This capability allows it to effectively integrate multiple data sources. In contrast, U-Net appears to rely more heavily on the specific channels it uses, with its performance being more sensitive to the quality and selection of those channels.

This difference in how the models handle feature importance reflects their underlying architectural differences. SegFormer, with its transformer-based design, is able to capture larger-scale patterns and dependencies within the data, giving it the flexibility to work with diverse and comprehensive input combinations. On the other hand, U-Net's pixel-wise segmentation approach tends to focus on more localized patterns, which can make it less adaptable when handling a variety of input channels.

Nevertheless, both models encountered certain challenges that influenced their performance, which we will now address in the following section.

781 6.2 Challenges Scenarios

782 The models in this study encountered several challenging scenarios related to the environmental
783 conditions of the test sites. These challenges, including riverbank areas affected by flooding,
784 geological variations, plowed land, and shadowed regions, directly influenced the models' ability to
785 detect landslides accurately. Spectral analysis of each scenario (Figure 14) reveals distinct patterns
786 that explain the models' performance limitations, while Figure 10 displays the corresponding mapping
787 results with recurrent error patterns.

788 *a) Riverbank Areas Post-Flooding*

789 The May 2023 flood events caused spectral signatures in riverbank areas that partially overlap with
790 those of landslides (Figure 14a). A comparison of pixels from non-landslide riverbank areas and those
791 affected by landslides reveals that while the spectral signatures are similar, there are noticeable
792 differences, particularly in the CGR RGB channels. These differences, although subtle, are
793 appreciable enough to help differentiate the two types of areas. Additionally, the lower slope values
794 in riverbank areas, due to their gentler topography along watercourses, and reduced vegetation change
795 indicators (NDVI Δ channels) set them apart from genuine slope failures. Post-flood sedimentation
796 and bank erosion generate surface disturbances that resemble landslide scars, especially where
797 floodwaters have stripped vegetation and redistributed sediments. Despite these similarities, these
798 features were not sufficiently weighted during training on the Casola Valsenio dataset, leading to
799 false positives in flood-affected environments. However, the distinction in the CGR RGB channels,
800 along with the differences in slope and NIR values, played a crucial role in limiting false negatives
801 (FN) along riverbanks, particularly in the best models (Caso 7, S7, and U7), despite the error still
802 being present.

803 *b) Geological Differences (FMA vs FAA)*

804 The geological difference between Marnoso-Arenacea (FMA) and Blue Clays (FAA) proved
805 challenging. Pixels from landslides in both lithologies were compared, with notable differences in the
806 CGR (Red) and CGR (NIR) channels, which are critical for landslide detection (Figure 14b). FAA
807 regions exhibit lighter colors in both pre- and post-event imagery, indicating typical erosion and
808 sediment redistribution in badlands (see Figure 10). This aligns with the diluvial processes observed
809 in the Blue Clays region. Training the models solely on FMA lithology, such as Casola Valsenio,
810 rendered them ineffective when applied to FAA regions, where the spectral characteristics differ
811 significantly. This highlights the limited transferability of models trained on a single lithology.

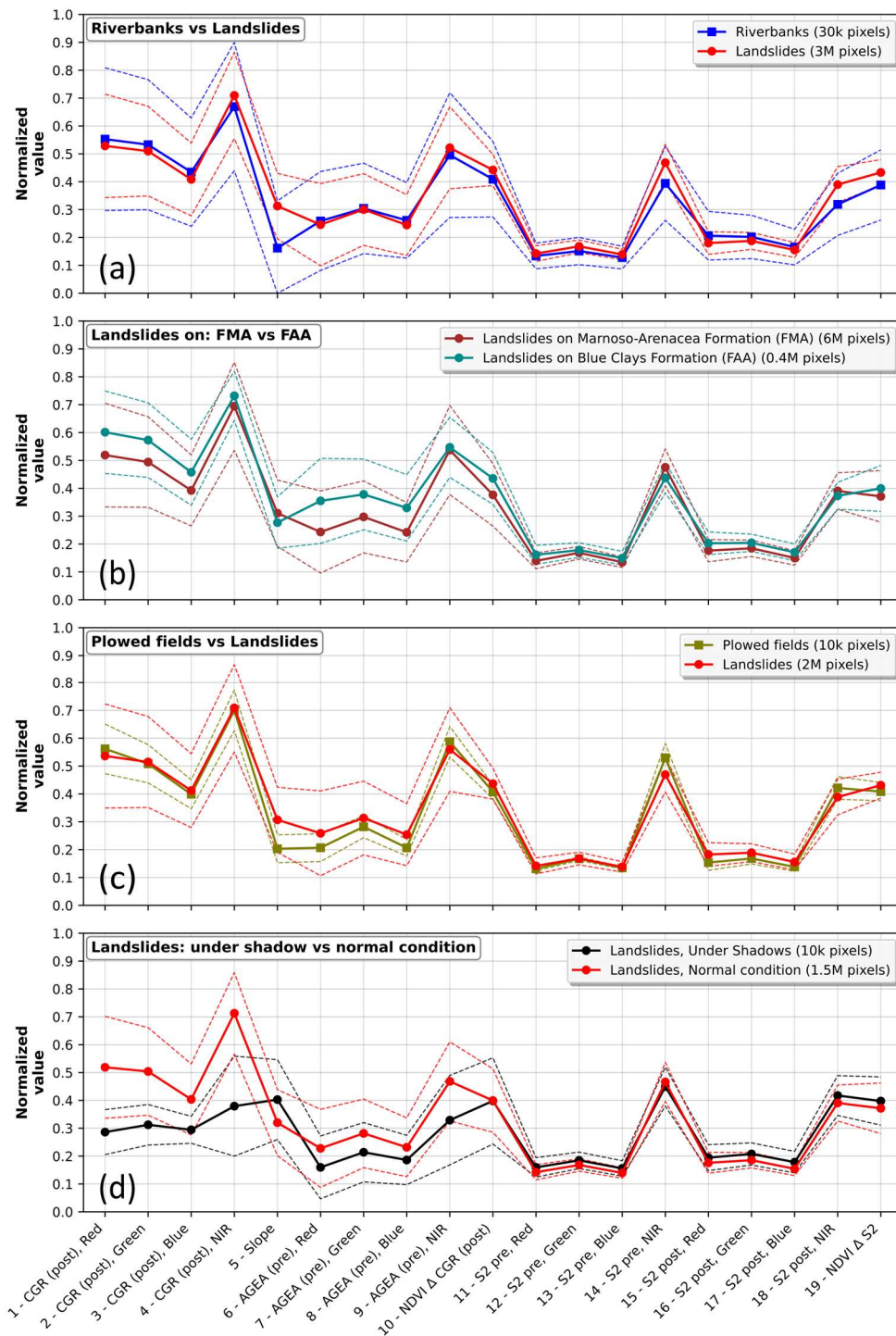
812 *c) Plowed Fields (Modigliana)*

813 Spectral signatures from landslide-affected areas in Modigliana were compared with those from
814 plowed fields outside the landslide zones (Figure 14c). The similarities in spectral signatures,
815 particularly in the CGR (Blue) and CGR (Green) channels, led to false positives (yellow polygons)
816 in Figure 10. Agricultural fields, when plowed, share similar characteristics with disturbed landslides,
817 especially when considering the differences in pre-event imagery. These differences are more
818 pronounced in the Slope and pre-event channels. However, due to the relatively low importance

819 assigned to these channels during training on the Casola Valsenio dataset, the models struggled to
820 differentiate these fields from landslides, resulting in confusion and false positives in plowed fields.

821 *d) Shadow Effects (Predappio)*

822 Shadowed regions, particularly in Predappio, made it nearly impossible for the models to classify
823 these areas as landslides. The spectral analysis (Figure 14d) highlights much darker values in
824 shadowed areas in the CGR (post-event) imagery, which complicates the recognition of landslides.
825 This is due to the drastically reduced reflectance in shadowed pixels, which prevents the models from
826 detecting landslides in these regions. Fortunately, this issue is isolated to a small portion of Predappio.
827 The reduced illumination and altered color profiles in shadowed areas hinder the model's ability to
828 detect the land movement, causing false negatives in these regions, as shown in Figure 10.



829

830 **Figure 14.** Spectral analysis of different land types and landslide detection. (a) Comparison between
831 riverbank areas affected by flooding and landslides. The blue line represents the spectral
832 characteristics of buildings along the river, while the red line corresponds to landslides. (b) Spectral
833 differences between landslides occurring on Marnoso-Arenacea Formation (FMA) and Blue Clays
834 (FAA), showing the significant spectral dissimilarity. (c) Comparison between spectral signatures of
835 plowed fields and landslides. The green line represents plowed fields, while the red line corresponds
836 to landslides, with noticeable overlap in spectral characteristics. (d) Impact of shadowed regions on
837 landslide detection, comparing normal conditions (red) with shadowed regions (black). The graph
838 reveals how shadows significantly reduce the ability of the models to detect landslides.

839 The analysis of various challenging scenarios, including flood-affected riverbanks, geological
840 variations, plowed land, and shadowed regions, reveals a consistent pattern: while identifiable
841 spectral discriminants exist for each challenge, the models' training on a single geological and
842 environmental domain (Casola Valsenio, FMA lithology) limited their ability to generalize across
843 varying conditions. The failure to adapt to different geological contexts, land use patterns, flood-
844 affected areas, and illumination conditions highlights the importance of domain-adaptive approaches
845 and more geographically comprehensive training datasets for robust landslide detection across
846 diverse landscapes.

847 To further investigate the role of geology in model generalization, an additional analysis was
848 conducted to assess the impact of including lithology among the input layers. For this, the lithological
849 layer was categorized into 8 distinct classes, with each class corresponding to a unique Unit ID (as
850 detailed in Table 2). These classes were numerically encoded from 1 to 8, representing different
851 lithological units. We compared the performance of Case S3 (RGB + slope) with the same
852 configuration plus the lithology layer. However, this did not yield significant improvements in
853 Casola, Predappio, Modigliana, or Brisighella MA, as these areas are predominantly characterized by
854 the Marnoso-Arenacea Formation ("Unit 7"), which was already well represented in the training
855 dataset (Figure 2a). As a result, the addition of the lithology layer provided redundant information.
856 However, when the same model was applied to Brisighella FAA, where the lithology is dominated
857 by Blue Clays ("Unit 1"), the model failed to detect any landslides, resulting in a very low F1-score
858 of 0.02. This underscores a key limitation: simply adding lithology as a static input layer is
859 insufficient for ensuring generalization. If the training dataset is not lithologically balanced, the model
860 may reinforce existing biases, associating landslide occurrence primarily with "Unit 7" and failing to
861 detect events in "Unit 1".

862 This finding highlights a fundamental issue in landslide detection models: effectively incorporating
863 lithological information requires more than just adding it as an input feature. The training dataset
864 itself must include landslide polygons from a representative range of lithological settings. When the
865 model was explicitly trained and tested on Brisighella FAA using the same configuration (Case S3:
866 RGB + slope), the F1-score increased from 0.32 to 0.41, confirming that geological consistency in
867 the training data significantly improves the model's ability to detect landslides in previously
868 underrepresented domains.

869 Although the F1-scores and Intersection over Union (IoU) achieved in this study may appear modest
870 compared to generic machine-learning benchmarks (e.g., Ghorbanzadeh et al., 2021, 2022; Prakash
871 et al., 2021; Meena et al., 2023), it is important to ground these results within the context of landslide
872 detection and segmentation literature, where lower scores are often observed. Several factors
873 contribute to these relatively low F1-scores, including class imbalance between landslide and non-
874 landslide areas, the spatial heterogeneity of landslides, uncertainties in reference inventories, and the
875 influence of image resolution and labeling quality. Landslide mapping, unlike conventional image
876 classification tasks, requires precise delineation of irregular and complex shapes across highly
877 variable terrain, which makes it more challenging to achieve high scores.

878 Our dataset was specifically designed for emergency response scenarios, where a small portion of the
879 affected area (Casola Valsenio) is mapped and the model is applied to the rest. This led to a train-
880 validation-test ratio of 12% for training, 3% for validation, and 85% for testing. The limited training
881 and validation sets, combined with the complex nature of landslide mapping, contribute to modest

882 F1-scores. Such a trade-off between accuracy and generalization is common in remote sensing, where
883 high-resolution, fine-grained feature extraction is both critical and difficult to achieve at a large scale.
884 Nevertheless, the potential of these automated approaches in emergency scenarios is considerable.
885 By rapidly generating landslide maps of comparable quality to expert-drawn products, these methods
886 could significantly accelerate initial response efforts. Reducing the need for time-consuming manual
887 digitization could save weeks of work, which is crucial during crisis events (Berti et al., 2025). The
888 automatically generated maps may serve as a robust initial product, enabling practitioners to focus on
889 refinement and validation rather than starting from scratch, ultimately delivering high-quality final
890 maps in a fraction of the time required for full manual mapping.

891

892 6.3 Identification of At-Risk Buildings

893 A further consideration in our study is the use of automated landslide maps to identify damage and
894 at-risk structures. Following the May 2023 disaster, the Emergency Commission determined that all
895 buildings located within landslide boundaries or within a 20-meter buffer were considered at risk and
896 potentially subject to relocation. To evaluate the effectiveness of our automated mapping products
897 for this task, we compared the buildings identified automatically with those derived from manual
898 mapping.

899 The F1-scores for the models indicate that "U2", "U4", "S6", and "U7" are the most accurate CNN
900 outputs, identifying on average 528 out of 654 buildings at risk (~0.78 F1-score). These results
901 demonstrate the models' ability to estimate the spatial extent of the phenomenon, even when using
902 Sentinel-2 imagery at 10 m resolution (U2). However, all models produced both false positives
903 (incorrectly flagged buildings) and false negatives (missed buildings), underscoring the need for
904 manual verification when high accuracy is required. Even a small number of misclassified structures,
905 especially inhabited buildings, can have serious consequences in emergency situations, thus
906 highlighting the importance of improving model performance for more effective response efforts.

907 Improving the landslide mapping itself is crucial for improving the accuracy of at-risk building
908 identification. As discussed in Section 7, refining mapping methodologies will lead to more reliable
909 and consistent results in future disaster scenarios.

910

911 6.4 Future Research Directions

912 To overcome the limitations identified in this study, future research should focus on key areas. A
913 primary challenge is optimizing data acquisition to improve map quality, including minimizing
914 shadow effects by scheduling imagery collection around noon, especially during winter months.
915 Expanding training datasets to include more plowed field examples will help models better
916 differentiate agricultural lands from landslides, reducing false positives. Further refinement could
917 include detailed land-use layers or slope filters to improve accuracy.

918 Improving data resolution is another important direction. While this study used 2-meter resolution
919 data, higher-resolution datasets (e.g., 20 cm) would provide more detailed information, enabling
920 better detection. Leveraging higher computational power would be necessary to process these datasets
921 and enhance model performance. In terms of approach, future research could explore advanced
922 techniques such as Multiscale Feature Pyramid Networks (FPN), which allow models to process

multiple scales simultaneously, Graph Neural Networks (GNNs) for capturing complex spatial relationships in geospatial contexts, and Neural Architecture Search (NAS) to automatically identify the best network architecture, improving robustness and generalization (Lin et al., 2017; Kipf & Welling, 2017; Zoph & Le, 2017).

Lastly, addressing the issue of geologically diverse training datasets remains crucial. Training models on datasets with limited geological diversity can hinder generalization. Expanding datasets to include various lithologies and geomorphological features will enhance model robustness. As deep learning models evolve, it is essential to prioritize the collection of diverse, high-quality data, as even the most advanced models cannot replace the need for comprehensive datasets. Future research should focus on improving data acquisition, enhancing model generalization, and refining validation techniques to further strengthen AI-based mapping in disaster management.

934

935 **7. Conclusion**

This study investigated the potential of automated landslide mapping to support rapid emergency response following the extreme meteorological events of May 2023 in Emilia-Romagna, Italy. Using a deep learning approach, we trained and tested two semantic segmentation models, U-Net and SegFormer, on high-resolution aerial imagery, Sentinel-2 data, NDVI change maps, and slope data. The training was conducted solely in the municipality of Casola Valsenio, while model performance was assessed on three additional municipalities (Brisighella, Modigliana, Predappio), chosen for their geological settings and significant landslide occurrence.

In the training area of Casola Valsenio, both U-Net and SegFormer achieved high performance, with consistent results across different input combinations. The models showed no significant variation in performance based on the input layers, as long as high-resolution post-event imagery (CGR) was included. However, when applied to external regions such as Brisighella, Modigliana, and Predappio, the models experienced a decrease in generalization. Despite this, the automated mapping still successfully identified the majority of landslides, demonstrating its utility in emergency scenarios.

In areas like Brisighella FAA (Blue Clays), where relevant lithologies were underrepresented in the training set, the models struggled to generalize effectively. This highlights the importance of incorporating a diverse range of lithologies in the training data to ensure robust model generalization across different geologically complex regions.

The choice of model architecture, whether U-Net or SegFormer, had minimal impact on performance, with both models showing similar results. Performance decreased with less detailed inputs, like Sentinel imagery alone, and improved with more rich inputs, reinforcing that data quality, not architecture, drives performance.

In emergency contexts, where external data for validation may not be available, expert judgment becomes essential for selecting the most reliable maps when ground truth data is lacking. While automated maps offer rapid assessments, they often require expert adjustments to refine outputs offering a balanced approach between speed and accuracy. The correct and thoughtful integration of AI-based systems into civil protection protocols represents a critical step forward, ensuring that these technologies complement existing procedures and improve overall response effectiveness.

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Disclosures and declarations

Declaration of AI and AI-assisted technologies in the writing process

During the preparation of this work the authors used ChatGPT 4 (chat.openai.com) to enhance the grammar and syntax, as well as to refine the sentence structure. All the content is original, and no concepts, ideas, or interpretations were produced by this tool. After using this tool, the authors reviewed and edited the content as needed and take full responsibility for the content of the publication.

Code and data availability

The dataset used in this study is openly available at <https://doi.org/10.5281/zenodo.13742643>.
The code developed for the analyses is available from the authors upon reasonable request.

Author contribution

NDS performed the data analysis, developed the scripts, and wrote the manuscript.
GC contributed to the interpretation of the results and supervised the work.
DE contributed to the methodology and algorithmic framework.
ELP contributed to the methodology and algorithmic framework.
AC supervised the interpretation of the results.
MB conceived the study, supervised the analysis and manuscript preparation, and provided critical review and guidance throughout the research process.

Competing interests

The authors declare they have no financial interests.

References

- Abrahams, E., Snow, T., Siegfried, M. R., & Pérez, F. (2024). A Concise Tiling Strategy for Preserving Spatial Context in Earth Observation Imagery. *arXiv preprint arXiv:2404.10927*.
- Amatya, P., Scheip, C., Déprez, A., et al. (2023). Learnings from rapid response efforts to remotely detect landslides triggered by the August 2021 Nippes earthquake and Tropical Storm Grace in Haiti. *Nat Hazards*, 118, 2337-2375. <https://doi.org/10.1007/s11069-023-06096-6>
- Amy, L. A., & Talling, P. J. (2006). Anatomy of turbidites and linked debrites based on long distance (120 × 30 km) bed correlation, Marnoso Arenacea Formation, Northern Apennines, Italy. *Sedimentology*, 53(1), 161-212. <https://doi.org/10.1111/j.1365-3091.2005.00756.x>
- Ban, Y., Zhang, P., Nascetti, A. et al. Near Real-Time Wildfire Progression Monitoring with Sentinel-1 SAR Time Series and Deep Learning. *Sci Rep* 10, 1322 (2020). <https://doi.org/10.1038/s41598-019-56967-x>
- Berti, M., Pizziolo, M., Scaroni, M. et al. Testing NDVI and U-Net for automated mapping of multiple-occurrence regional landslide events using satellite and aerial multispectral data (Casola Valsenio, Emilia-Romagna, Northern Apennines, Italy). *Landslides* (2026). <https://doi.org/10.1007/s10346-025-02671-z>
- Berti, M., Pizziolo, M., Scaroni, M., Generali, M., Critelli, V., Mulas, M., Tondo, M., Lelli, F., Fabbiani, C., Ronchetti, F., Ciccarese, G., Dal Seno, N., Ioriatti, E., Rani, R., Zuccarini, A., Simonelli, T., & Corsini, A. (2025). RER2023: The landslide inventory dataset of the May 2023 Emilia-Romagna meteorological event. *Earth System Science Data*, 17(3), 1055-1074. <https://doi.org/10.5194/essd-17-1055-2025>
- Blaschke, T. (2010). Object based image analysis for remote sensing. *ISPRS Journal of Photogrammetry and Remote Sensing*, 65(1), 2–1. <https://doi.org/10.1016/j.isprsjprs.2009.06.004>
- Brath, A., Casagli, N., Marani, M., Mercogliano, P., & Motta, R. (2023). Report by the Technical-Scientific Commission established by Regional Council Resolution No. 984/2023 and Managerial Determination No. 14641/2023 to analyze the extreme meteorological events of May 2023. Technical-Scientific Commission Report. Retrieved June 15, 2024, from <https://www.regione.emilia-romagna.it/alluvione/>
- Casagli, N., Cigna, F., Bianchini, S., et al. (2016). Landslide mapping and monitoring by using radar and optical remote sensing: examples from the EC-FP7 project SAFER. *Remote Sens Appl Soc Environ*, 4, 92-108.
- Catani, F., Tofani, V., & Raspini, F. (2020). Use of remote sensing for landslide studies. *Progress in Landslide Science*, 75-102.
- Chen, Z., Zhang, Y., Ouyang, C., Zhang, F., & Ma, J. (2018). Automated landslides detection for mountain cities using multi-temporal remote sensing imagery. *Sensors (Basel)*, 18, 821. <https://doi.org/10.3390/s18030821>

1034 Chen, X., Liu, M., Li, D., Jia, J., Yang, A., Zheng, W., & Yin, L. (2023). Conv-trans dual network
1035 for landslide detection of multi-channel optical remote sensing images. *Frontiers in Earth*
1036 *Science*, 11, 1182145. <https://doi.org/10.3389/feart.2023.1182145>

1037 Cruden, D.M., Varnes, D.J.: *Landslide Types and Processes*, Transportation Research Board, U.S.
1038 National Academy of Sciences, Special Report, 247, 36-75, 1996

1039 Dal Seno, N., Evangelista, D., Piccolomini, E., Berti, M. Comparative analysis of conventional and
1040 machine learning techniques for rainfall threshold evaluation under complex geological
1041 conditions. *Landslides* (2024). <https://doi.org/10.1007/s10346-024-02336-3>

1042 Dosovitskiy, A., et al. (2020). An Image is Worth 16x16 Words: Transformers for Image
1043 Recognition at Scale. arXiv preprint arXiv:2010.11929.
1044 <https://doi.org/10.48550/arXiv.2010.11929>

1045 Ferrario, M. F., & Livio, F. (2023). Rapid Mapping of Landslides Induced by Heavy Rainfall in
1046 the Emilia-Romagna (Italy) Region in May 2023. *Remote Sens (Basel)*, 16, 122.
1047 <https://doi.org/10.3390/rs16010122>

1048 Foraci, R., Tesini, M. S., Nanni, S., Antolini, G., & Pavan, V. (2023). L'inquadramento meteo e
1049 idrologico degli eventi. *Ecoscienza, ARPAE Emilia-Romagna*, anno XIV, 5, 20-24.

1050 Ghorbanzadeh, O., Blaschke, T., Gholamnia, K., Meena, S. R., Tiede, D., & Aryal, J. (2019).
1051 Evaluation of different machine learning methods and deep-learning Convolutional Neural
1052 Networks for landslide detection. *Remote Sensing*, 11(2), 196.
1053 <https://doi.org/10.3390/rs11020196>

1054 Ghorbanzadeh, O., Xu, Y., Ghamisi, P., Kopp, M., & Kreil, D. (2022). Landslide4Sense: Reference
1055 Benchmark Data and Deep Learning Models for Landslide Detection. *IEEE Transactions on*
1056 *Geoscience and Remote Sensing*, 60. <https://doi.org/10.48550/arXiv.2206.00515>

1057 Ghorbanzadeh, O., Shahabi, H., Crivellari, A., Homayouni, S., Blaschke, T., & Ghamisi, P. (2022).
1058 Landslide detection using deep learning and object-based image analysis. *Landslides*, 19(4),
1059 929-939. <https://doi.org/10.1007/s10346-021-01843-x>

1060 Ghorbanzadeh O., Khalil Gholamnia & Pedram Ghamisi (2023) The application of ResU-net and
1061 OBIA for landslide detection from multi-temporal Sentinel-2 images, *Big Earth Data*, 7:4,
1062 961-985, DOI: 10.1080/20964471.2022.2031544.

1063 Guzzetti, F., Mondini, A. C., Cardinali, M., et al. (2012). Landslide inventory maps: New tools for
1064 an old problem. *Earth Sci Rev*, 112, 42-66. <https://doi.org/10.1016/j.earscirev.2012.02.001>.

1065 Holbling, D., Eisank, C., Albrecht, F., et al. (2017). Comparing manual and semi-automated
1066 landslide mapping based on optical satellite images from different sensors. *Geosciences*.
1067 <https://doi.org/10.3390/geosciences7020037>.

1068 Isensee, F., Jaeger, P.F., Kohl, S.A.A. *et al.* nnU-Net: a self-configuring method for deep learning-
1069 based biomedical image segmentation. *Nat Methods* **18**, 203-211 (2021).
1070 <https://doi.org/10.1038/s41592-020-01008-z>

Iverson, R. M., George, D. L., Allstadt, K., et al. (2015). Landslide mobility and hazards: implications of the 2014 Oso disaster. *Earth Planet Sci Lett*, 412, 197-208. <https://doi.org/10.1016/j.epsl.2014.12.020>

LeCun, Y., Bengio, Y. & Hinton, G. Deep learning. *Nature* **521**, 436-444 (2015). <https://doi.org/10.1038/nature14539>

Lin, T.-Y., Dollár, P., Girshick, R., He, K., & Hariharan, B. (2017). Feature Pyramid Networks for Object Detection. *CVPR 2017*. <https://arxiv.org/abs/1612.03144>

Liu, X., Zhao, C., Zhang, Q., Lu, Z., Li, Z., Yang, C., Zhu, W., Liu-Zeng, J., Chen, L., & Liu, C. (2021). Integration of Sentinel-1 and ALOS/PALSAR-2 SAR datasets for mapping active landslides along the Jinsha River corridor, China. *Engineering Geology*, 284, 106033. <https://doi.org/10.1016/j.enggeo.2021.106033>

Lu, W., Hu, Y., Zhang, Z., Chen, L., & Liu, C. (2023). A dual-encoder U-Net for landslide detection using Sentinel-2 and DEM data. *Landslides*, 20(8), 1975-1987. <https://doi.org/10.1007/s10346-023-02089-5>

Ji, S. P., Yu, D. W., Shen, C. Y., Li, W. L., & Xu, Q. (2020). Landslide detection from an open satellite imagery and digital elevation model dataset using attention boosted convolutional neural networks. *Landslides*, 17, 1337-1352. <https://doi.org/10.1007/s10346-020-01353-2>.

Kingma, D. P., & Ba, J. (2017). Adam: A Method for Stochastic Optimization. arXiv preprint. <https://doi.org/10.48550/arXiv.1412.6980>

Kipf, T. N., & Welling, M. (2017). Semi-Supervised Classification with Graph Convolutional Networks. *ICLR 2017*. <https://arxiv.org/abs/1609.02907>

Nava, L.; Bhuyan, K.; Meena, S.R.; Monserrat, O.; Catani, F. Rapid Mapping of Landslides on SAR Data by Attention U-Net. *Remote Sens.* 2022, 14, 1449. <https://doi.org/10.3390/rs14061449>

Notti, D., Cignetti, M., Godone, D., & Giordan, D. (2023). Semi-automatic mapping of shallow landslides using free Sentinel-2 images and Google Earth Engine. *Natural Hazards and Earth System Sciences*, 23(9), 2625-2648. <https://doi.org/10.5194/nhess-23-2625-2023>

Novellino, A., Pennington, C., Leeming, K. et al. Mapping landslides from space: A review. *Landslides* **21**, 1041-1052 (2024). <https://doi.org/10.1007/s10346-024-02215-x>

Martha, T. R., Kerle, N., Jetten, V., van Westen, C. J., & Kumar, K. V. (2010). Characterising spectral, spatial and morphometric properties of landslides for semi-automatic detection using object-oriented methods. *Geomorphology*, 116, 24–36. <https://doi.org/10.1016/j.geomorph.2009.10.004>

Meena, S. R., Ghorbanzadeh, O., & van Westen, C. J., et al. (2021). Rapid mapping of landslides in the Western Ghats (India) triggered by 2018 extreme monsoon rainfall using a deep learning approach. *Landslides*, 18, 1937-1950. <https://doi.org/10.1007/s10346-020-01602-4>

Meena, S. R., Soares, L. P., & Grohmann, C. H., et al. (2022). Landslide detection in the Himalayas using machine learning algorithms and U-Net. *Landslides*, 19, 1209-1229. <https://doi.org/10.1007/s10346-022-01861-3>

1110 Meena, S. R., Nava, L., Bhuyan, K., Puliero, S., Soares, L. P., Dias, H. C., Floris, M., and Catani, F.
 1111 (2023). HR-GLDD: a globally distributed dataset using generalized deep learning (DL) for
 1112 rapid landslide mapping on high-resolution (HR) satellite imagery, *Earth Syst. Sci. Data*, 15,
 1113 3283-3298, <https://doi.org/10.5194/essd-15-3283-2023>

1114 Mondini, A. C., Santangelo, M., Rocchetti, M., Rossetto, E., Manconi, A., & Monserrat, O. (2019).
 1115 Sentinel-1 SAR Amplitude Imagery for Rapid Landslide Detection. *Remote Sensing*, 11(7),
 1116 760. <https://doi.org/10.3390/rs11070760>

1117 Prakash, N., Manconi, A., & Loew, S. (2020). Mapping Landslides on EO Data: Performance of Deep
 1118 Learning Models vs. Traditional Machine Learning Models. *Remote Sensing*, 12(3), 346.
 1119 <https://doi.org/10.3390/rs12030346>

1120 Prakash, N., Manconi, A. & Loew, S. A new strategy to map landslides with a generalized
 1121 convolutional neural network. *Sci Rep* 11, 9722 (2021). [https://doi.org/10.1038/s41598-021-](https://doi.org/10.1038/s41598-021-89015-8)
 1122 [89015-8](https://doi.org/10.1038/s41598-021-89015-8)

1123 N. Prakash and A. Manconi (2021). Rapid Mapping of Landslides Triggered by the Storm Alex,
 1124 October 2020. 2021 IEEE International Geoscience and Remote Sensing Symposium
 1125 IGARSS, Brussels, Belgium, 2021, pp. 1808-1811.
 1126 <https://doi.org/10.1109/IGARSS47720.2021.9553321>

1127 Paszke, A., Gross, S., Massa, F., Lerer, A., Bradbury, J., Chanan, G., ... & Chintala, S. (2019).
 1128 PyTorch: An Imperative Style, High-Performance Deep Learning Library. In *Advances in*
 1129 *Neural Information Processing Systems* (pp. 8026-8037).

1130 Pizziolo, M., Generali, M., & Scaroni, M. (2023). In Appennino un numero di frane mai
 1131 riscontrato prima. *Ecoscienza, ARPAE Emilia-Romagna*, anno XIV, 5, 31-33.

1132 Pizziolo, M., Berti, M., Scaroni, M., Generali, M., Critelli, V., Mulas, M., Tondo, M., Lelli, F.,
 1133 Fabbiani, C., Ronchetti, F., Ciccarese, G., Dal Seno, N., Ioriatti, E., Rani, R., Zuccarini, A.,
 1134 Simonelli, T., & Corsini, A. (2024). RER2023: the landslide inventory dataset of the May
 1135 2023 Emilia-Romagna event - Version 1 (Version 1) [Data set]. Zenodo.
 1136 <https://doi.org/10.5281/zenodo.13742643>

1137 Prakash, N., Manconi, A. & Loew, S. A new strategy to map landslides with a generalized
 1138 convolutional neural network. *Sci Rep* 11, 9722 (2021).
 1139 <https://doi.org/10.1038/s41598-021-89015-8>

1140 Prechelt, L. (1998). Early stopping-but when? In *Neural Networks: Tricks of the Trade* (pp. 55-
 1141 69). Springer, Berlin, Heidelberg.

1142 Rezatofighi, H., Tsoi, N., Gwak, J., Sadeghian, A., Reid, I., & Savarese, S. (2019). Generalized
 1143 intersection over union: A metric and a loss for bounding box regression. In *Proceedings of*
 1144 *the IEEE/CVF Conference on Computer Vision and Pattern Recognition* (pp. 658-666).

1145 Ronneberger, O., Fischer, P., & Brox, T. (2015). U-Net: Convolutional Networks for Biomedical
 1146 Image Segmentation. *International Conference on Medical Image Computing and*
 1147 *Computer-Assisted Intervention*, 234-241.

1148 Sameen, M. I., & Pradhan, B. (2019). Landslide detection using residual networks and the fusion
 1149 of spectral and topographic information. *IEEE Access*, 7, 114363-114373.
 1150 <https://doi.org/10.1109/access.2019.2935761>

1151 Tang, X., Tu, Z., Wang, Y., Liu, M., Li, D., & Fan, X. (2022). Automatic Detection of Coseismic
 1152 Landslides Using a New Transformer Method. *Remote Sens*, 14, 2884.
 1153 <https://doi.org/10.3390/rs1412f2884>

1154 Tharwat, A. (2020). Classification assessment methods. *Applied Computing and Informatics*,
 1155 17(1), 168-192.

1156 Wasowski, J., & Bovenga, F. (2014). Investigating landslides and unstable slopes with satellite Multi
 1157 Temporal Interferometry: Current issues and future perspectives. *Engineering Geology*, 174,
 1158 103-138. <https://doi.org/10.1016/j.enggeo.2014.03.003>

1159 Wolf, T., Debut, L., Sanh, V., Chaumond, J., Delangue, C., Moi, A., ... & Rush, A. M. (2020).
 1160 Transformers: State-of-the-art Natural Language Processing. In *Proceedings of the 2020*
 1161 *Conference on Empirical Methods in Natural Language Processing: System Demonstrations*
 1162 (pp. 38-45).

1163 Wu, L., Liu, R., Ju, N., Zhang, A., Gou, J., He, G., & Lei, Y. (2024). Landslide mapping based on a
 1164 hybrid CNN-transformer network and deep transfer learning using remote sensing images
 1165 with topographic and spectral features. *International Journal of Applied Earth Observation*
 1166 *and Geoinformation*, 126, 103612. <https://doi.org/10.1016/j.jag.2023.103612>

1167 Xie, E., Wang, W., Yu, Z., Anandkumar, A., Alvarez, J. M., & Luo, P. (2021). SegFormer:
 1168 Simple and Efficient Design for Semantic Segmentation with Transformers. *Advances in*
 1169 *Neural Information Processing Systems*, 34, 12077-12090.

1170 Xu, Y., Ouyang, C., Xu, Q. *et al.* CAS Landslide Dataset: A Large-Scale and Multisensor Dataset for
 1171 Deep Learning-Based Landslide Detection. *Sci Data* **11**, 12 (2024).
 1172 <https://doi.org/10.1038/s41597-023-02847-z>

1173 Yang, W., Wang, Y., Sun, S. *et al.* Using Sentinel-2 time series to detect slope movement before the
 1174 Jinsha River landslide. *Landslides* **16**, 1313-1324 (2019).
 1175 <https://doi.org/10.1007/s10346-019-01178-8>

1176 Ye CM et al (2019) Landslide detection of hyperspectral remote sensing data based on deep learning
 1177 with constrains. *IEEE Journal of Selected Topics in Applied Earth Observations and Remote*
 1178 *Sensing* 12:50475060. <https://doi.org/10.1109/JSTARS.2019.2951725>.

1179 Zhang, Z., Liu, Q., & Wang, Y. (2018). Road Extraction by Deep Residual U-Net, in *IEEE*
 1180 *Geoscience and Remote Sensing Letters*, vol. 15, no. 5, pp. 749-753, May 2018.
 1181 doi:10.1109/LGRS.2018.2802944.

1182 Zoph, B., & Le, Q. V. (2017). Neural Architecture Search with Reinforcement Learning. *ICLR 2017*.
 1183 <https://arxiv.org/abs/1611.01578>
 1184