



1 **PM_{2.5} Assimilation within JEDI for NOAA's Regional Air** 2 **Quality Model (AQMv7): Application to the September 2020** 3 **Western U.S. Wildfires**

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19 **Abstract.** This paper describes efforts to establish aerosol data assimilation capabilities for NOAA's National Air
20 Quality Forecasting Capability (NAQFC), a regional online air quality modeling (AQM) system under NOAA's
21 Unified Forecast System (UFS), by assimilating measurements of fine particulate matter (PM_{2.5}, particles with
22 diameters less than 2.5 µm). PM_{2.5} assimilation is developed within the Joint Effort for Data assimilation Integration
23 (JEDI) framework and tested using its 3D-Var data assimilation (DA) component. The PM_{2.5} observation operator is
24 constructed by combining newly developed PM_{2.5} transformation recipes in the JEDI Variable Derivation
25 Repository (VADER) with a general spatial interpolation operator in the Unified Forward Operator (UFO).

26 Cycled DA and forecast experiments were conducted from 1 to 21 September 2020, during a period of Western U.S.
27 wildfires, to assess the impact of assimilating PM_{2.5} observations from the AirNow and PurpleAir networks. The
28 control and analysis variables include individual aerosol species, with background error standard deviations
29 generated by scaling their respective background values. Prognostic variables such as aerosol particle number and
30 total particulate surface area are updated accordingly following each analysis update. All DA experiments use a
31 3-hourly cycling interval, with PM_{2.5} observations assimilated every 3 hours. The control experiment uses the same
32 configuration but without any data assimilation. Results show that assimilating either AirNow or PurpleAir PM_{2.5}
33 data reduces 1–24 h forecast errors in terms of mean absolute error (MAE) and root mean square error (RMSE)



34 compared to the control run over CONUS. Forecast skill, measured using the Critical Success Index (CSI) for $\text{PM}_{2.5}$
35 thresholds of 5, 12, and $35 \mu\text{g}/\text{m}^3$, also improves. AirNow observations have a greater impact overall, while
36 PurpleAir shows its strongest impact over Nevada, northern Utah, Colorado, and southwestern New
37 Mexico—regions with persistent underpredictions in the control run at forecast hour 1. Overall, the assimilation of
38 PurpleAir observations in addition to AirNow data leads to a slight reduction in 3–24 h MAE.

39 1 Introduction

40 Particulate matter with an aerodynamic diameter of 2.5 micrometers or smaller ($\text{PM}_{2.5}$) is a major contributor to
41 poor air quality in the United States, posing significant risks to public health and the environment, and contributing
42 to substantial loss of life. Over the past few decades, poor air quality in the U.S. has contributed to over 100,000
43 premature deaths annually, far exceeding fatalities from all other weather-related causes combined, which average
44 around 500 per year (Huang et al., 2025). Given its public health significance, $\text{PM}_{2.5}$ is one of the primary pollutants
45 used in calculating the Air Quality Index (AQI)—a standardized system designed to communicate daily air pollution
46 levels to the public. Elevated $\text{PM}_{2.5}$ concentrations frequently result in "unhealthy" AQI ratings, triggering health
47 advisories and public warnings.

48 $\text{PM}_{2.5}$ in the United States originates from a range of both anthropogenic and natural sources. Anthropogenic
49 sources include agricultural activities and combustion processes, such as emissions from motor vehicles, power
50 plants, industrial facilities, and residential heating systems. Among natural sources, wildfires are a particularly
51 significant contributor, especially in the western United States, where their frequency and intensity have escalated
52 dramatically over the past two decades (Wen and Burke, 2021). According to the U.S. Environmental Protection
53 Agency (EPA), wildfires account for approximately 15% to 30% of total $\text{PM}_{2.5}$ emissions nationwide (EPA, 2017).
54 While national seasonal averages of $\text{PM}_{2.5}$ have generally declined, summer $\text{PM}_{2.5}$ concentrations in the western
55 U.S. have remained persistently high, primarily due to wildfire smoke (O'Dell et al., 2019). In addition to degrading
56 air quality, wildfires have caused widespread property loss. Since 2005, more than 99,500 homes, businesses, and
57 other structures have been destroyed by wildfire-related events
58 (<https://headwaterseconomics.org/natural-hazards/structures-destroyed-by-wildfire>, last access on June 30, 2025),
59 underscoring the urgent need for more effective strategies in air quality monitoring, forecasting, and wildfire
60 management.

61 The National Oceanic and Atmospheric Administration (NOAA) has developed an advanced regional Air Quality
62 Modeling (AQM) prediction system within the Unified Forecast System (UFS) framework to enhance the accuracy
63 of air quality forecasts across the United States, particularly during wildfire events (Huang et al. 2025). The National
64 Air Quality Forecast Capability (NAQFC), operated by NOAA's National Weather Service (NWS), has been
65 providing operational air quality forecast guidance for over 20 years, with continuous inclusion of new capabilities.
66 Under NAQFC, the AQM version 7 was implemented and became operational on May 14, 2024. A key innovation
67 in this system is the integration of the Real-time Aerosol and fire behavior Visual Estimator (RAVE) — a high



68 spatiotemporal resolution, satellite-derived wildfire product — which enables a more accurate representation of
69 wildfire emissions. The system also features online coupling of atmospheric and chemical models, allowing
70 dynamic interactions between meteorology and atmospheric chemistry. This integration improves the representation
71 of emissions and ensures real-time feedback of meteorological fields that influence chemical transformations and the
72 transport of pollutants in the atmosphere. The UFS-AQM online system has consistently shown improved
73 performance in simulating major wildfire events, including the significant wildfires in the northwestern coastal
74 regions of the U.S. in September 2020, and widespread smoke transport from Canadian wildfires in the summer of
75 2023. This system was officially implemented on May 14, 2024 as NOAA's operational air quality prediction system
76 (AQMV7), replacing the previous offline-coupled the Global Forecast System using the Finite Volume Cube-Sphere
77 dynamical core (GFS-FV3) version 15 with the Community Multiscale Air Quality modeling system (CMAQv5.0.2)
78 modeling system. (Chen et al. 2021).

79 PM_{2.5} data assimilation (DA) has proven effective in reducing errors in air quality forecasts (e.g., Pagowski et al.
80 2010, 2012; Schwartz et al. 2012; Wu et al. 2015; Robichaud 2017; Lee et al. 2021; Chen et al. 2022, Ha 2022;
81 Vogel et al. 2025, among others). Pagowski et al. (2010) demonstrated that fine aerosol forecasts benefit from
82 AirNow PM_{2.5} DA, showing improved verification scores for a period of at least 24 hours. Schwartz et al. (2012)
83 found that assimilating AirNow PM_{2.5} observations significantly improved surface PM_{2.5} forecasts over the
84 CONUS compared to forecasts without DA. Wu et al. (2015) reported that incorporating ground-based PM_{2.5}
85 observations notably enhanced 24-hour forecasts during a severe pollution episode in Shanghai. Similarly, Chen et
86 al. (2022) showed that assimilating multi-source PM_{2.5} data significantly improved WRF-Chem PM_{2.5} forecasts
87 with benefits lasting up to 48 hours. Lee et al. (2021) highlighted the effectiveness of assimilating ground in-situ
88 surface PM_{2.5} observations in improving the short-term PM_{2.5} predictions in Northeast Asia.

89 Many operational regional air quality prediction systems around the world use some form of data assimilation to
90 initialize the forecasts. These approaches vary in complexity, ranging from simple optimal interpolation to full
91 variational or ensemble Kalman filter methods (e.g. Robichaud et al. 2016; Wei et al. 2024; Colette et al. 2024). In
92 NOAA's current regional air quality model (AQMV) operations, aerosol and chemical initial conditions are
93 "warm-started" using 6-hour forecasts from the previous model cycle. The implementation of an aerosol data
94 assimilation system can further enhance short-term air quality forecasts by providing more accurate spatial analyses
95 of initial aerosol distributions.

96 To establish aerosol data assimilation capabilities for NOAA's regional operational AQMV system, we employ the
97 Joint Effort for Data assimilation Integration (JEDI) (Trémolet and Auligné, 2020). JEDI is a flexible, agnostic, and
98 modern data assimilation system applicable to a wide range of forecasting systems (e.g. Liu et al. 2023; Huang et al.
99 2023; Sluka, 2024). JEDI offers a platform that supports efficient scientific development and facilitates the transition
100 from research to operations. As part of a broader strategic shift, NOAA and partner agencies are transitioning their
101 data assimilation systems to JEDI, opening the door for rapid integration of new scientific advancements, greater



consistency across modeling systems, and enhanced collaboration across research communities and operational centers.

This study aims to develop an initial aerosol analysis capability for the AQM system by assimilating $\text{PM}_{2.5}$ observations using the JEDI 3D-Var framework. Low-cost PurpleAir data are valuable for real-time air quality monitoring and are displayed in the AirNow Fire and Smoke Map (<https://fire.airnow.gov/>, last access: July 19 2025). However, their impact on numerical air quality prediction has not been thoroughly studied. In addition to AirNow $\text{PM}_{2.5}$ measurements, this study also evaluates the impact of assimilating PurpleAir observations.

The paper is organized as follows: section 2 provides a description of Methodology including the NOAA's AQM system, 3D-Var approach, and JEDI $\text{PM}_{2.5}$ assimilation. Experimental setup is presented in section 3 including case description, AQM configuration, AirNow and PurpleAir $\text{PM}_{2.5}$ observations and background errors setup. Results are described in section 4. A summary and conclusion are presented in the final section.

2 Methodology

2.1 AQMv7 overview

The NOAA's regional operational AQMv7 system was developed through the online coupling of the Finite-Volume version 3 (FV3) dynamical core -based atmospheric model (Black et al 2021) with the EPA's Community Multiscale Air Quality (CMAQ) model v5.2.0 within the UFS framework (Huang et al., 2025). In this UFS-AQM online system, CMAQ is treated as an atmospheric chemistry column model to simulate atmospheric chemistry reactions that govern concentrations of chemical species including gas- and aerosol-phase species. The transport terms of chemical species are handled by the FV3 dynamical core in the same way as other physics tracers (Huang et al., 2025). Aerosol module version 6 (AERO6) (Zhang et al. 2018) is utilized by CMAQ to simulate aerosol processes.

The AQMv7 system is configured over the North American domain with a grid-spacing of 13 km and 65 vertical levels, extending up to 0.2 hPa. In total, AERO6 simulates 76 aerosol-related variables. Additional information about the UFS-AQM online system can be found in Huang et al. (2025). In this research, the model configuration is the same as the operational AQMv7 setup except for running over the CONUS domain with a 3 hourly cycling interval.

2.2 $\text{PM}_{2.5}$ assimilation within JEDI 3D-Var

In the JEDI framework, a series of components are provided to create a flexible, comprehensive data assimilation system. The JEDI three-dimensional variational (3D-Var) component is used to assimilate $\text{PM}_{2.5}$ for AQMv7. The 3D-Var method is chosen for its operational feasibility, primarily due to its low computational cost and the fact that it does not require an ensemble prediction system.



In practice, a 3D-Var data assimilation system typically uses an incremental approach to minimize a quadratic cost function which is defined in terms of the analysis increment δx relative to the guess state x_g :

$$J(\delta x) = \frac{1}{2} (\delta x - \delta x_g) B^{-1} (\delta x - \delta x_g)^T + \frac{1}{2} (\mathbf{H}[\delta x] - d) R^{-1} (\mathbf{H}[\delta x] - d)^T \quad (1)$$

Where:

- $\delta x_g = x_b - x_g$ is the guess state departure from background state x_b , which is usually taken from a previous short-term forecast.
- \mathbf{H} is the linearized observation operator of nonlinear observation operator \mathbf{H} .
- \mathbf{B} and \mathbf{R} are the background and observation error covariance matrices, respectively.
- d is the innovation vector, defined as:

$$d = y - H(x_g) \quad (2)$$

with y representing the observation vector.

Once the increment δx is obtained, the analysis state x^a is reconstructed as:

$$x^a = x_g + \delta x \quad (3)$$

2.2.1 PM_{2.5} observation operator

In AQMv7, the modal approach taken in the CMAQ model represents aerosol particle size distributions as the superposition of three lognormal modes: Aitken (I), accumulation (J), and coarse (K). It predicts only three integral properties of the size distribution for each mode: the total particle number concentration, the total surface area concentration, and the total mass concentration of the individual chemical components.

The total PM_{2.5} concentration is calculated as a weighted sum of the individual aerosol concentration across these three modes:

$$\text{PM}_{2.5} = \text{ATOTI} \cdot \text{PM}_{25\text{AT}} + \text{ATOTJ} \cdot \text{PM}_{25\text{AC}} + \text{ATOTK} \cdot \text{PM}_{25\text{CO}} \quad (4)$$

Here, ATOTI, ATOTJ, and ATOTK represent the total aerosol mass concentrations in the Aitken, accumulation, and coarse modes, respectively. For example, ATOTI is the combined mass of 14 prognostic aerosol variables in the Aitken mode from the AERO6 aerosol module. Similarly, ATOTJ and ATOTK are the aggregated mass concentrations of 49 and 7 aerosol variables in the accumulation and coarse modes, respectively. PM_{25AT}, PM_{25AC}, and PM_{25CO} are mass scaling factors for the three modes that vary by location and time. The aerosol variables within the same mode share the same mass scaling factor.

The PM_{2.5} observation operator is constructed by combining PM_{2.5} transformation recipes in the JEDI Variable



Derivation Repository (VADER) with a general spatial interpolation operator in the Unified Forward Operator (UFO). VADER is responsible for transforming model variables using user-defined “recipes” to generate new variables in model space. For $PM_{2.5}$ assimilation, VADER computes $PM_{2.5}$ from individual aerosol species using model-specific transformation, specifically using the equation 4 for this application. Since $PM_{2.5}$ composition varies by model, these transforms are implemented within VADER to match the specific structure of the regional air quality model AQMv7. Once $PM_{2.5}$ is derived in model space, UFO applies a generic spatial interpolation operator to map the model-simulated values to the observation locations, enabling computation of the observed minus forecast values.

The input for the $PM_{2.5}$ transformation are mixing ratio of the 70 aerosol variables wrt dry air in unit ug/kg, the three mass scaling factors in the three modes, and dry air density for unit conversion. The output product is the $PM_{2.5}$ in unit ug/m^3 . It is noted that a recipe that uses temperature, surface pressure, and delta pressure to derive the dry air density in case the dry air density is not found in the input variable list into VADER.

The JEDI/VADER $PM_{2.5}$ recipe provides nonlinear (NL), tangent linear (TL), and adjoint (AD) transforms of $PM_{2.5}$ that keeps the output products in the same grid space as the input variables. Hence, the generic interpolation operator in UFO is used to connect the model-derived 3D $PM_{2.5}$ fields with observed surface $PM_{2.5}$ measurements. This respects the JEDI paradigm of keeping the UFO part of the JEDI model independent.

2.2.2 Background error covariance modeling

In a 3D-Var system, the background error covariance (BEC) determines both the spatial spreading of information from observations and the magnitude of the analysis increments along with the observation error variance.

The background error covariance matrix B can be decomposed into a standard deviation matrix (Σ) and a correlation matrix (C), as follows:

$$B = \Sigma C \Sigma \quad (5)$$

The correlation matrix C is generally non-diagonal. Σ is a diagonal matrix, with the standard deviations of the background errors for each variable on the diagonal.

The error modeling of the correlation matrix and standard deviations usually apply to control variables. In the first implementation of aerosol data assimilation in JEDI for AQMv7, the control variables are defined as individual forecast aerosol variables, resulting in 70 control variables for AQMv7 with the AERO6 aerosol mechanism. The setup of background error standard deviation and correlation modeling will be described in Section 3: Experimental setup.

2.2.3 Minimization Algorithm (DRIPCG)



JEDI provides several minimization algorithm options. In this paper, we use the Derber–Rosati Inexact Preconditioned Conjugate Gradient (DRIPCG) algorithm (Derber and Rosati, 1989), as implemented in the JEDI's OOPS (Object-Oriented Prediction System) framework. DRIPCG has been extensively tested and is chosen here for stability and convergence efficiency.

3 Experimental setup

3.1 The September 2020 fire event

The wildfires of September 2020 ranked among the most intense in the U.S. in recent years. These fires produced dense smoke that initially moved westward over the Willamette Valley and eventually blanketed the broader region. As a result, air quality rapidly deteriorated to hazardous levels, marking one of the worst air quality periods in recent decades (Mass et al., 2021). Wildfire smoke originating from California, Oregon, and Washington was injected into the free troposphere and transported across the country by prevailing winds, leading to hazy conditions in several states. According to Li et al. (2021), from August to October 2020, wildfires in the western U.S. contributed 23% of surface $\text{PM}_{2.5}$ across the contiguous United States (CONUS), with higher contributions observed along the Pacific Coast (43%) and in the Mountain Region (42%). This study focuses on the peak fire activity occurring between September 1 and 21.

3.2 $\text{PM}_{2.5}$ observations

In this study, surface $\text{PM}_{2.5}$ observations were obtained from two sources: AirNow and PurpleAir observing networks. These datasets differ in sensor type, spatial coverage, and quality control (QC) requirements. AirNow provides regulatory-grade measurements from federal, state, and local monitoring stations, while PurpleAir is a low-cost, community-based network of air quality sensors. PurpleAir sensors are widely deployed by individuals and communities, providing real-time data on $\text{PM}_{2.5}$ concentrations as well as meteorological variables such as temperature, pressure, and relative humidity. Only the data reported from outdoor $\text{PM}_{2.5}$ sensors are used in this study. The PurpleAir data were available for registered users through the PurpleAir API (<https://community.purpleair.com/t/api-use-guidelines/1589>)

3.2.1 PurpleAir $\text{PM}_{2.5}$ quality control and correction

Quality control and correction of PurpleAir data followed the methodology described in Barkjohn et al. (2021). Readers are referred to that paper for further details. A correction is required because the PurpleAir raw data usually overestimate $\text{PM}_{2.5}$ concentrations under typical ambient and smoke-impacted conditions. The following quality control (QC) filters were applied to the raw PurpleAir $\text{PM}_{2.5}$ measurements:

- Reported $\text{PM}_{2.5}$ values from two Plantower sensors within the PurpleAir sensor (channels A and B) must be nonnegative.



- 221 • The PurpleAir sensor channel A and B consistency:
- 222 ○ Absolute difference $< 5 \mu\text{g}/\text{m}^3$, *or*
- 223 ○ Relative difference within 61%.
- 224 • $\text{PM}_{2.5}$ values must not exceed PM_{10} values.
- 225 • $\text{PM}_{2.5}$ values must be less than $3000 \mu\text{g}/\text{m}^3$ (upper threshold).
- 226 • Gross check of relative humidity with range 0-100%.

227 Only PurpleAir $\text{PM}_{2.5}$ measurements that passed all the above QC criteria were retained for subsequent correction.

228 3.2.2 PurpleAir $\text{PM}_{2.5}$ correction

229 Correction of PurpleAir $\text{PM}_{2.5}$ measurements was performed using a multiple linear regression model based on
 230 sensor-reported $\text{PM}_{2.5}$ (PA) and relative humidity (RH), following the correction formula proposed by Barkjohn et
 231 al. (2021):

$$232 \text{PM}_{2.5} = 0.524 \times \text{PA} - 0.0862 \times \text{RH} + 5.75 \quad (6)$$

233 We adopt the above equation because it was United States-wide valid by fitting data from September 2017 until
 234 January 2020. Though the above correction equation is originally for 24h averaged $\text{PM}_{2.5}$, a similar regression
 235 equation was derived from the September 2020 1h averaged $\text{PM}_{2.5}$ dataset:

$$236 \text{PM}_{2.5} = 0.508 \times \text{PA} - 0.0449 \times \text{RH} + 4.89 \quad (7)$$

237 The close similarity between the two equations supports the consistency and robustness of the correction method
 238 across datasets and time periods.

239 To reduce random sensor noise and improve comparability with the model resolution (~ 13 km), the corrected
 240 PurpleAir $\text{PM}_{2.5}$ data were spatially averaged onto a $0.1^\circ \times 0.1^\circ$ latitude–longitude grid. PurpleAir shows a good
 241 coverage of Washington, Oregon, California and Colorado, and more observations of Arizona, Utah, New Mexico,
 242 Texas.

243 3.2.3 Observation error assignment

244 Observation error standard deviations were assigned to each network:

- 245 • AirNow $\text{PM}_{2.5}$: 5% of observed value
- 246 • PurpleAir $\text{PM}_{2.5}$: 10% of observed value

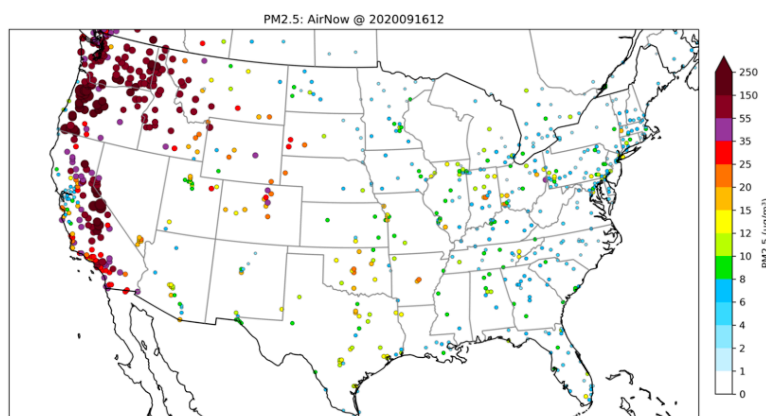
247 The values above are based on the EPA’s definition of acceptable measurement uncertainty, which specifies a 10%
 248 coefficient of variation for total precision. The AirNow $\text{PM}_{2.5}$ observation errors were set to 5% of the observed
 249 values. Park et al. (2022) also used a 5% error specification for assimilating $\text{PM}_{2.5}$ observations, though their study



250 focused on observation networks over China and Korea. For PurpleAir PM_{2.5} data, the observation errors were set
 251 to 10%, reflecting the higher likelihood of greater uncertainties associated with lower-cost sensors. For comparison,
 252 the default PM_{2.5} observation error configuration in Gridpoint Statistical Interpolation (GSI) (Pagowski et al., 2012;
 253 Wang et al., 2021) includes a measurement error modeled as $1.5 + 0.75\% \times \text{PM}_{2.5}$, along with a representativeness
 254 error component. At the current model resolution of 13 km, the error specification used in this study reduces the
 255 influence of large PM_{2.5} observations, particularly those exceeding approximately 55 $\mu\text{g}/\text{m}^3$ for AirNow monitors
 256 and 25 $\mu\text{g}/\text{m}^3$ for PurpleAir sensors.

257 Figure 1a–b shows the spatial distribution of AirNow and PurpleAir PM_{2.5} monitoring stations at 1200 UTC on
 258 September 16, 2020. Figures 1c–d display the time series of domain averaged PM_{2.5} values and station counts from
 259 the AirNow and PurpleAir networks, including matched stations between the two. PurpleAir sensors are especially
 260 concentrated in densely populated areas, leading to notable spatial variability in observation coverage during the
 261 September 2020 wildfire events. Coverage is particularly dense in urban regions of the western United States (e.g.
 262 California, Oregon, Washington, Utah, Arizona and Colorado), while rural and remote areas have significantly fewer
 263 sensors, for example, Nevada and North Dakota. The number of AirNow stations ranges from approximately 800 to
 264 900, while PurpleAir stations number between 1,160 and 1,300. Dropouts in the AirNow network lead to sudden
 265 decreases in station count and corresponding drops in the PM_{2.5} time series. In contrast, the PurpleAir network
 266 shows a general upward trend in station count, with no major data dropouts observed.

267 (a)



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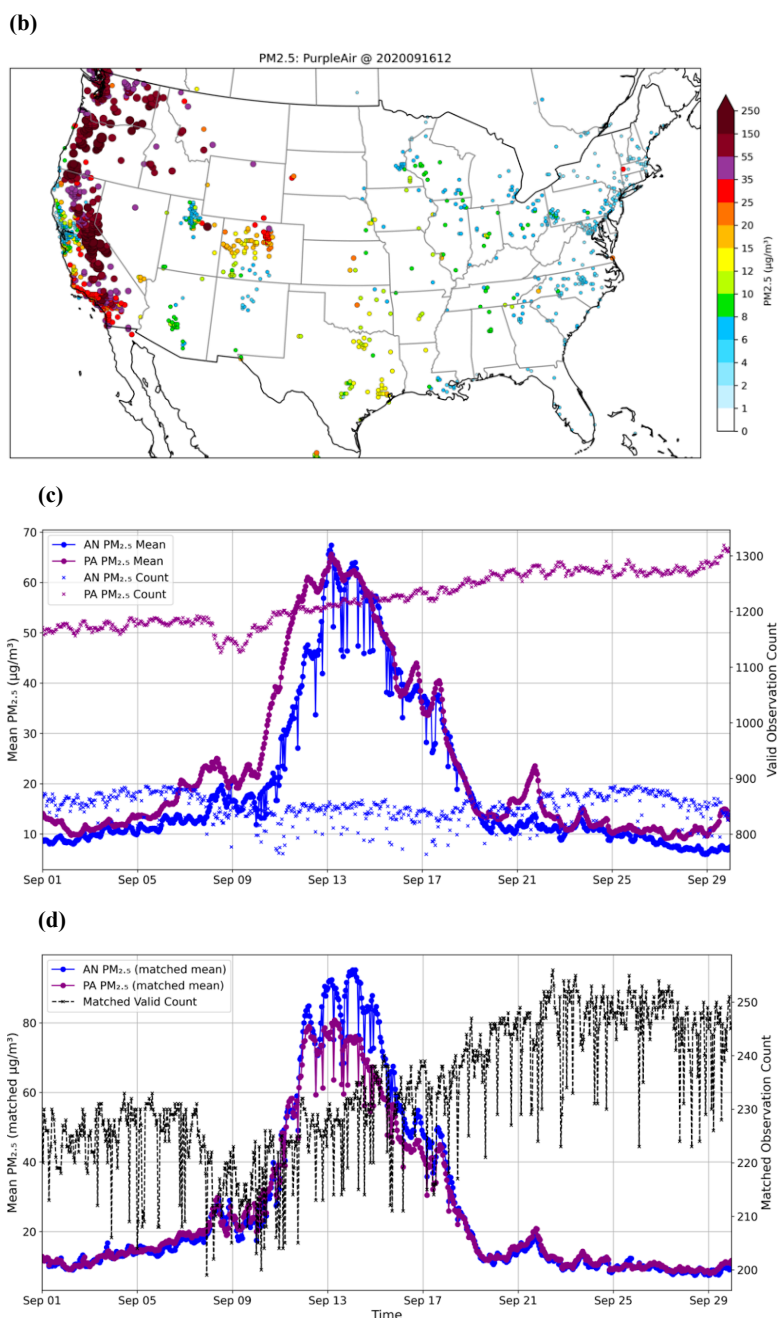


Figure 1. (a-b). Spatial distribution of AirNow(AN) and PurpleAir(PA) PM_{2.5} monitoring stations on 1200 UTC 16 September 2020. (c). Time series of domain averaged PM_{2.5} values and numbers from AirNow and PurpleAir observing networks. (d) Time series of domain averaged PM_{2.5} values and numbers for matched AirNow and PurpleAir stations.



285 3.3 Background error covariance

286 In this study, the background error standard deviation (Σ) for each control variable is constructed based on the
 287 background forecast; specifically, the error standard deviations of an aerosol variable are prescribed as proportional
 288 to its background values.

289 The proportional scaling factor s is approximately estimated by building a linear relationship between the $PM_{2.5}$
 290 standard error (Σ) and the background forecast $PM_{2.5}^{bkg}$ $PM_{2.5}$ concentrations:

$$291 \Sigma = s \cdot PM_{2.5}^{bkg} \quad (8)$$

292 The scaling factor s is subsequently applied to all $PM_{2.5}$ components, i.e., the 70 prognostic aerosol variables, to
 293 construct their error standard deviations.

294 This proportionality-based approach has also been adopted in the MOCAGE operational system (Colette et al.,
 295 2024), where background error standard deviations are similarly prescribed relative to background concentrations as
 296 a first-order approximation.

297 Tang et al. (2023) tested a similar method, in which the background $PM_{2.5}$ error variance (Σ^2) is first estimated using
 298 the Hollingsworth–Lönnberg method (Hollingsworth and Lönnberg, 1986). A linear relationship is then established
 299 between the estimated $PM_{2.5}$ standard error (Σ) and the background forecast $PM_{2.5}^{bkg}$.

300 Here we take the same idea but using an alternative approach to roughly estimate the background $PM_{2.5}$ forecast
 301 error variance (Σ^2). The background $PM_{2.5}$ error variance (Σ^2) is estimated using $PM_{2.5}$ innovation information d_b^o
 302 and observation error information \mathcal{E}^o defined in the subsection 2.2, specifically,

$$303 E(\mathcal{E}^b \mathcal{E}^{bT}) = E(d_b^o d_b^{oT}) - E(\mathcal{E}^o \mathcal{E}^{oT}) \quad (9)$$

304 Equation 9 is valid under the assumption that observation and background errors are uncorrelated. This assumption
 305 is reasonable when the innovation vector d_b^o is calculated using forecasts from a free-running model without any
 306 aerosol data assimilation.

307 In this study, short-term (e.g., 3-hour) $PM_{2.5}$ forecasts from a free run conducted during 1–21 September 2020 were
 308 used to compute the innovation vector d defined in Equation 2. This free run, referred to as the *control run*, is
 309 described in detail in the following section. Based on the innovations and observation errors defined in subsection
 310 3.2.3, which serve as inputs to Equation 9 for estimating background standard deviation error of $PM_{2.5}$, then a
 311 scaling factor s was estimated using Equation 8, with the background $PM_{2.5}$ standard deviation error and
 312 background values as inputs. This scaling factor was subsequently applied in all assimilation experiments presented

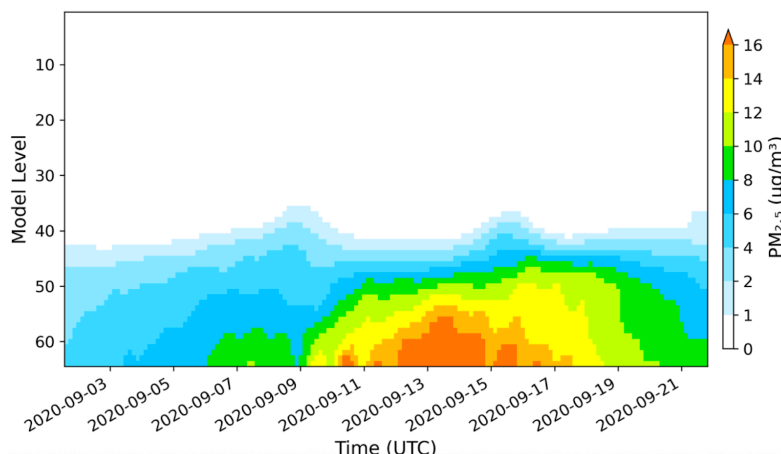


313 in this study.

314 This proportionality-based approach implicitly assumes that displacement errors in background don't dominate,
 315 focusing the assimilation process on correcting amplitude. It offers several benefits:

- 316 • It helps constrain analysis increments to physically meaningful regions. For example, it prevents the
 317 generation of sea salt aerosol increments over inland areas where no sea salt is present in the background.
 318 This is a problem that can occur when using GSI's height-dependent or latitude–height-dependent
 319 background error variance formulations, particularly when individual aerosol species are used as control
 320 variables.
- 321 • It introduces location- and time-dependent background error variance information, improving the realism of
 322 background error specification. Moreover, the aerosol variables that dominate background errors vary by
 323 location and assimilation cycle, rather than being consistently dominated by the same species when using
 324 constant static background error statistics. For example, organic and black carbon typically exhibit the
 325 largest errors in wildfire regions and downwind areas affected by smoke, whereas other regions may be
 326 dominated by non-organic aerosols.

327 An example of background error standard deviation in $PM_{2.5}$ space from a data assimilation run that assimilated
 328 both AirNow and PurpleAir $PM_{2.5}$ is shown in Figure 2. This figure is intended to illustrate the main difference to
 329 static constant background errors, though the actual errors used in the data assimilation experiments are the errors of
 330 the individual aerosol control variables. It is obvious that this approach produces dynamically location- and
 331 time-dependent varying error estimates that yields particularly large error variances during the peak fire events from
 332 10 to 20 September 2020.



333

334 **Figure 2.** Domain averaged $PM_{2.5}$ standard deviations for the data assimilation run that assimilated both AirNow
 335 and PurpleAir $PM_{2.5}$.



The background error correlation matrix C is modeled using a generic diffusion correlation operator designed for short length scales, as implemented in the System-Agnostic Background Error Representation (SABER) repository (Sluka, 2024). A horizontal cutoff scale of 100 km is applied, consistent with estimates derived from NMC statistics in previous GSI applications (Wang et al., 2021). For vertical correlations, this study uses a cutoff scale of 12 model levels, which helps confine the influence of surface $PM_{2.5}$ observations within the average daytime planetary boundary layer (PBL) height (~ 1450 m) and has demonstrated improved surface $PM_{2.5}$ prediction as will be discussed in Section 4.

3.4 Update of total particle number and surface area concentrations

After the aerosol mass concentration has been analyzed, total particle number concentration, total surface area concentration can be updated accordingly. For simplicity, it is assumed that the ratio of the particle number concentration to total particulate volume within each mode (I, J, K) remains the same as in the background. Total particulate volume is used instead of mass mixing ratio because it is proportional to the particle number concentration (see Eq. 3 in Binkowski and Roselle, 2003). A similar assumption was adopted by Li (2013) to update number concentrations for the WRF-Chem model.

The number of particles is updated using the following relation:

$$N_a = N_b / V_b \times V_a \quad (10)$$

Where:

- N_a and N_b are the number of particles in the analysis and background, respectively, within each mode.
- V_a and V_b are the total particulate volumes in the analysis and background, respectively, within the same mode.

The total particulate volume (V_a or V_b) within each mode is calculated by dividing the mass concentration of each aerosol variable by its corresponding density in that mode, and then summing the results. This updating approach implicitly assumes that changes in volume across the three modes are driven solely by variations in particle number, rather than shifts in the aerosol size distribution. The total particulate surface area within each mode is then updated using the same volume ratio, i.e., V_a/V_b (Eq.10) multiplied by the background surface area.

In preparatory work for this study, six-hourly cycling experiments (Wang et al., 2025) have shown that updating these variables is crucial for improving AQMv7 performance. In contrast, previous work using GSI with earlier developmental versions of AQM did not update these variables, primarily because those model versions were less advanced than the current operational AQMv7. As a result, there was still significant room for improving prediction



365 skills.

366 **3.5 Experiments**

367 Table 1 provides a description of the experiments. Four experiments were conducted to evaluate the performance of
368 JEDI/AQM PM_{2.5} DA. The first experiment is a control run (CTR), in which meteorological initial and boundary
369 conditions are updated every 3 hours, while chemical and aerosol fields are carried over from the 3-hour forecast of
370 the previous cycle. The other three experiments incorporate data assimilation: DA_AN, DA_PA, and DA_ANPA,
371 which assimilate AirNow PM_{2.5} only, PurpleAir PM_{2.5} only, and both AirNow and PurpleAir PM_{2.5} observations,
372 respectively.

373 Like the CTR experiment, all DA experiments are conducted as 3-hourly cycling runs, with PM_{2.5} observations
374 assimilated every 3 hours. 24-hour forecasts are initialized four times daily at 0000 UTC, 0600 UTC, 1200 UTC,
375 and 1800 UTC. The experimental period spans from 1200 UTC on September 1 to 1800 UTC on September 21,
376 2020.

377 **Table 1. Descriptions of the experiments.**

Experiment	Data Assimilation	PM _{2.5} Observations Assimilated	Aerosol Fields
CTR	No	None	Carried over from previous cycle's 3-hour forecast
DA_AN	Yes	AirNow PM _{2.5} only	Updated by Assimilation
DA_PA	Yes	PurpleAir PM _{2.5} only	Updated by Assimilation
DA_ANPA	Yes	AirNow + PurpleAir PM _{2.5}	Updated by Assimilation

378

379 **4 Results**

380 **4.1. Results from all cycles**

381 This section provides an overview of the impact of DA on PM_{2.5} forecasts. A total of 80 forecasts—initialized four



times daily from 0000 UTC on September 2 to 1800 UTC on September 21, 2020—are used to evaluate model performance. AirNow PM_{2.5} observations are used to verify the forecast performance. Forecast errors are assessed using bias, mean absolute error (MAE), and root mean square error (RMSE). Forecast performance is evaluated using box plots and performance diagrams. The box-and-whisker plots illustrate the distribution, spread, and central tendency of forecast errors, while the performance diagrams highlight forecast skill (e.g., Critical Success Index, CSI). Time series of PM_{2.5} at various forecast hours are presented to examine the temporal evolution of forecast performance. Additionally, spatial distributions of PM_{2.5} including observations, forecasts, forecast errors, and forecast differences are analyzed to evaluate the spatial impact of data assimilation on PM_{2.5} predictions. AirNow PM_{2.5} observations are used as reference to evaluate forecast skills.

Figure 3 presents the bias, mean absolute error (MAE), and root mean square error (RMSE) for the 1–24 h forecast of domain-averaged PM_{2.5}. Domain averages are computed over EPA Regions 1–10, which include all states in the mainland United States. The detailed description of EPA regions can be found on EPA webpage: <https://www.epa.gov/aboutepa/regional-and-geographic-offices#regional>, last access on 11 July 2025. Overall, all data assimilation experiments show improved forecast skill compared to the control run. The added value of assimilating PurpleAir PM_{2.5} data alongside AirNow observations is evident in the consistent MAE reduction (Fig. 3b). Its impact on RMSE is also positive, though relatively small.

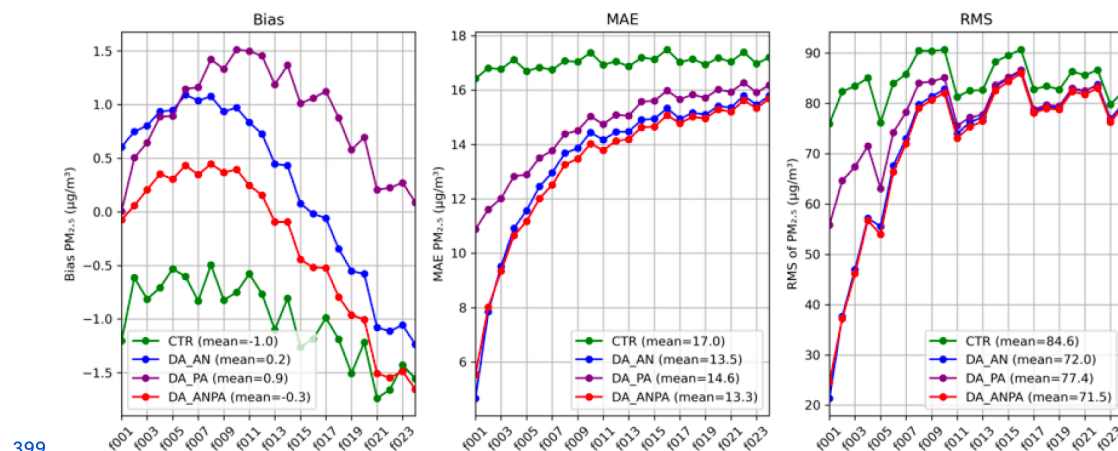


Figure 3. PM_{2.5} forecast errors for 1–24 h lead times based on 80 forecasts initialized four times daily during September 2–21, 2020. Domain-averaged over EPA Regions 1–10.

(a) Bias, (b) Mean Absolute Error (MAE), (c) Root Mean Square Error (RMSE).

Figure 4 shows box-and-whisker plots of PM_{2.5} forecast bias. Across all forecast hours, the interquartile range (IQR)—represented by the height of the boxes—is consistently smaller for the DA experiments compared to the control run. This indicates reduced forecast error spread between the 25th and 75th percentiles and suggests more consistent forecasts in the DA experiments. Although the median forecast bias in the control run is sometimes closer to zero, the DA_ANPA experiment performs comparably in terms of central tendency while showing clear



improvements in reducing the mean forecast bias, as also reflected in Figure 3a. Among the DA experiments, DA_AN and DA_ANPA show the most consistent improvement at 24-hour lead times, with DA_ANPA slightly outperforming others during the early forecast hours (e.g., hour 1 to 12). This suggests that assimilating PurpleAir observations in addition to AirNow helps reduce bias and brings the forecasts closer to observed $\text{PM}_{2.5}$ values in the short term.

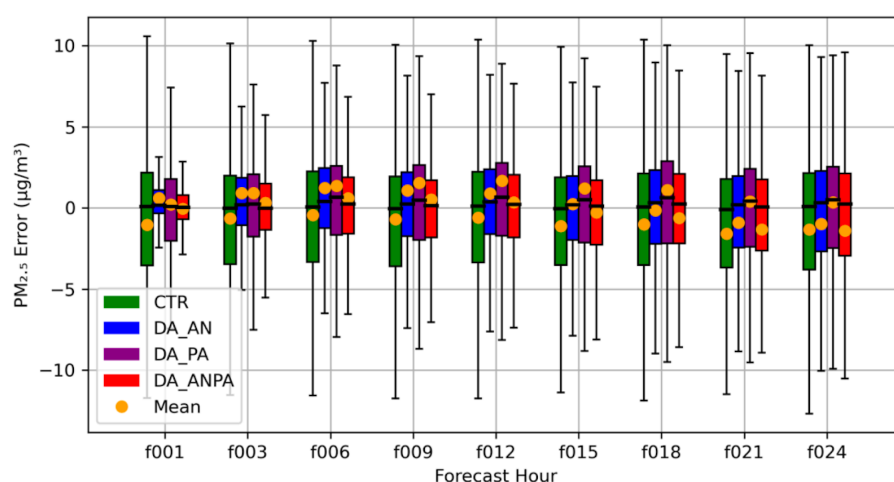
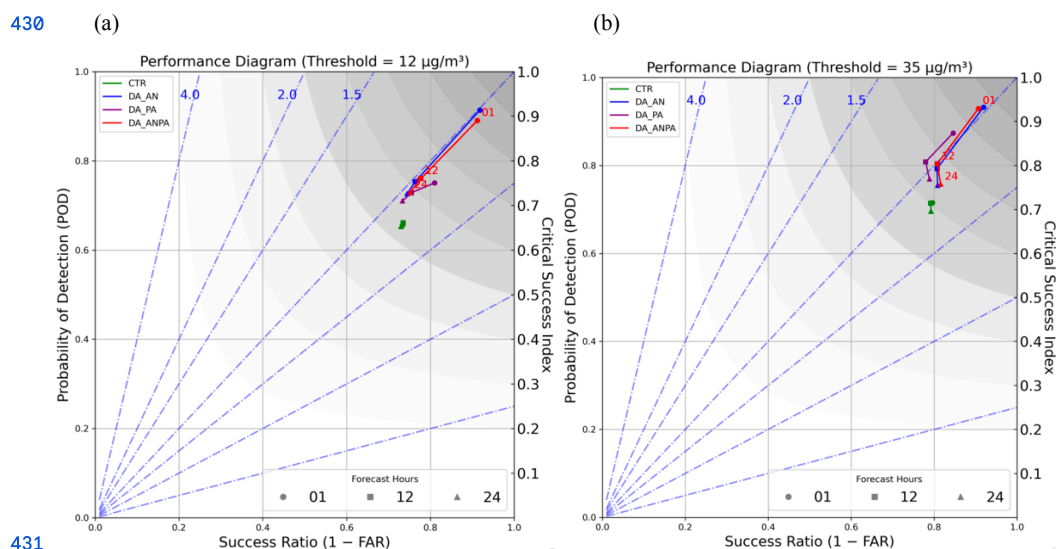


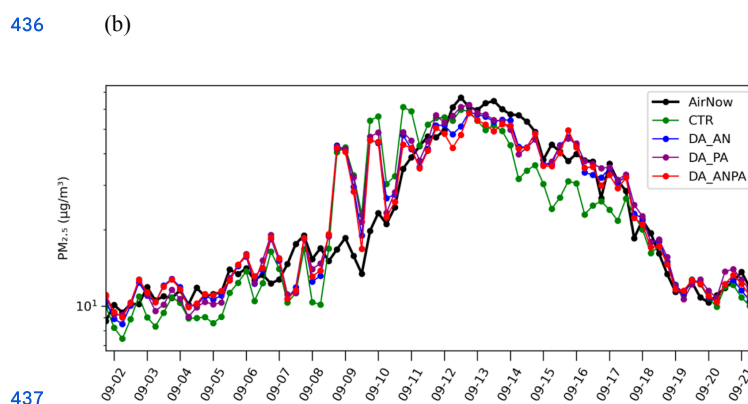
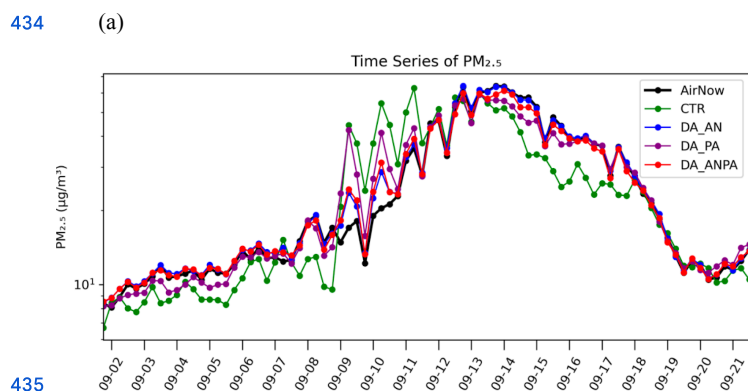
Figure 4. Box-and-whisker plot of $\text{PM}_{2.5}$ forecast bias. Bottom edge = Q1 (25th percentile); Top edge = Q3 (75th percentile); Height = Interquartile Range ($\text{IQR} = \text{Q3} - \text{Q1}$); Horizontal line inside box: The median (50th percentile); Whiskers: Extend to the min and max values within $1.5 \times \text{IQR}$ from Q1 and Q3.

Figure 5 displays performance diagrams of $\text{PM}_{2.5}$ forecast at forecast hours 1, 12, and 24 with $\text{PM}_{2.5}$ threshold of $12 \mu\text{g}/\text{m}^3$ and $35 \mu\text{g}/\text{m}^3$. Performance diagrams show consistent improvements in CSI scores across all forecast hours for all DA experiments, with DA_AN and DA_ANPA outperforming the DA_PA experiment. The performance with $\text{PM}_{2.5}$ threshold of $5 \mu\text{g}/\text{m}^3$ (figure not shown) is similar to that of $12 \mu\text{g}/\text{m}^3$.

Figure 6 presents shows time series of $\text{PM}_{2.5}$ averaged over EPA Regions 1–10 at forecast hours 1, 12, and 24, respectively. Consistent with the evaluations in Figures 3 and 4, all DA experiments generally improve $\text{PM}_{2.5}$ forecasts. Notably, all DA experiments help correct underpredictions during September 2–9 and 14–17. In addition, the substantial overprediction during September 10–13 observed in the control run, largely due to inaccurate fire emissions, is partially mitigated by the DA experiments. Among the DA configurations, DA_AN and DA_ANPA show comparable performance and both outperform DA_PA.

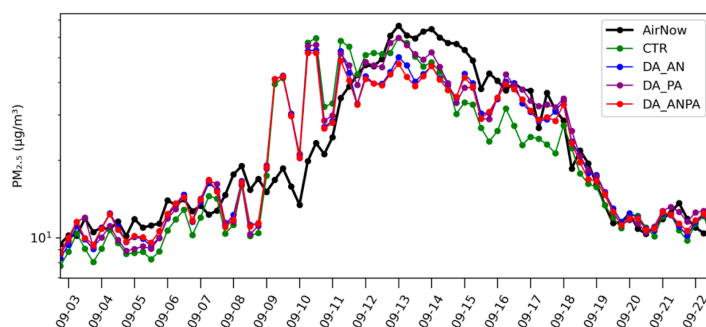


432 **Figure 5.** (a). Performance diagram for forecast hours 1, 12, and 24 with a $\text{PM}_{2.5}$ threshold of $12 \mu\text{g}/\text{m}^3$. (b) Same as
 433 (a), but using a threshold of $35 \mu\text{g}/\text{m}^3$.





438 (c)



439

440 **Figure 6.** Time series of $PM_{2.5}$ averaged over EPA Regions 1–10 for (a) forecast hour 1, (b) forecast hour 12, and
 441 (c) forecast hour 24. The y-axis is shown on a logarithmic scale.

442 While we have investigated the impact of DA on $PM_{2.5}$ forecasts in terms of temporal evolution, it is also important
 443 to examine the spatial distribution of forecast fields, associated errors, and how DA influences these spatial patterns.

444 Figure 7 presents the spatial distribution of temporally averaged $PM_{2.5}$ forecasts at forecast hour 1, based on 80
 445 forecasts initialized four times daily (0000 UTC, 0600 UTC, 1200 UTC, and 1800 UTC) from September 2 to 21.
 446 Figure 7a shows the $PM_{2.5}$ field from the control run (shaded), overlaid with AirNow observations. The effects of
 447 wildfire events are clearly seen across California, Oregon, and Washington—where the fires occurred—as well as in
 448 downstream regions impacted by smoke advection and transport.

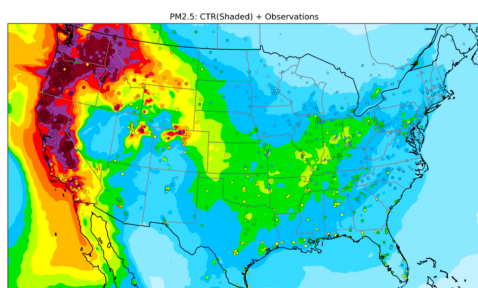
449 Forecast errors in the control run are evident in Figure 6a but are more clearly highlighted in Figure 7b, which shows
 450 the difference between the control run and AirNow observations. Significant overpredictions appear along the
 451 California coast, as well as in parts of the Midwest and Northeast U.S., including Tennessee, Kentucky, West
 452 Virginia, and Virginia, which are approximately represented by EPA regions 1, 2, 3, 5 and 7. Conversely, notable
 453 underpredictions are found over Colorado, New Mexico, much of Texas and Oklahoma, and several Gulf Coast
 454 states which are in EPA regions 4 and 6.

455 Compared to the control run, both DA_AN (Fig. 7c-d) and DA_PA (Fig. 7e-f) show similar spatial correction
 456 patterns across California, Oregon, and Washington, particularly in reducing overpredictions along the California
 457 coast. They also produce comparable large-scale adjustments across the Northeast, Midwest, and Southern U.S.,
 458 with error patterns (Fig. 7d and 7f) largely opposite in sign to those in the CTR–AirNow difference (Fig. 7b). This
 459 suggests that both DA experiments effectively mitigate the control run’s over- and underpredictions.

460 However, the magnitude of correction is generally smaller in DA_PA than in DA_AN. Notably, DA_PA shows its
 461 strongest impact over Nevada, northern Utah, Colorado, and southwestern New Mexico, helping to alleviate the
 462 underpredictions in these regions—similar to improvements seen in DA_ANPA.

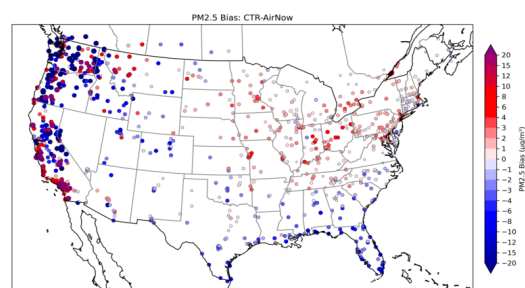


463 (a) CTR

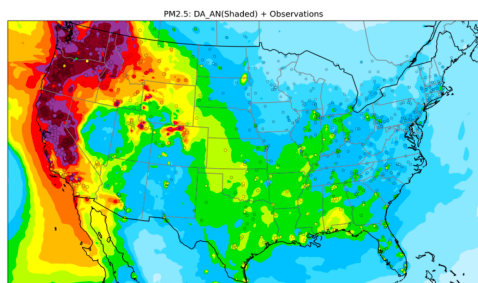


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(b) CTR-AirNow

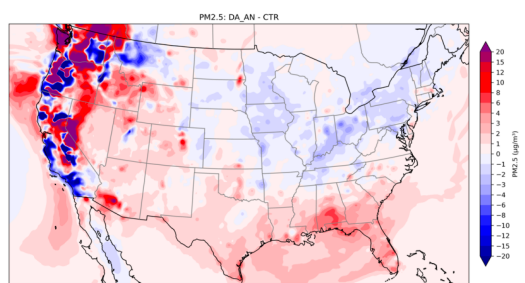


465 (c) DA_AN

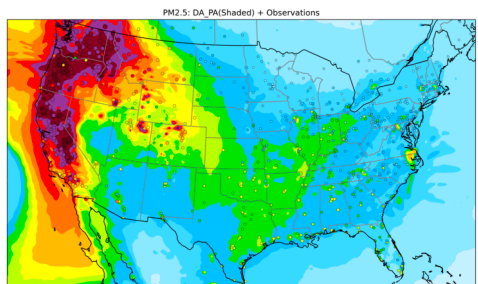


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(d) DA_AN-CTR

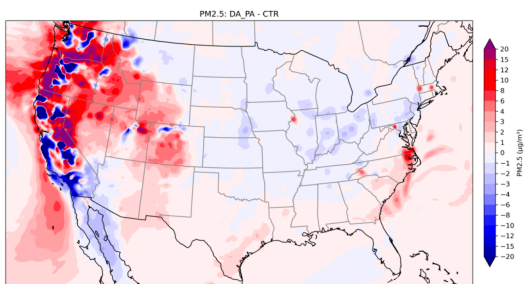


467 (e) DA_PA

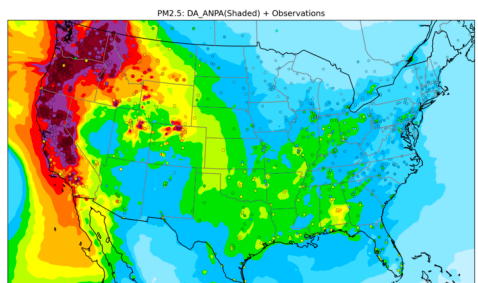


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(f) DA_PA-CTR

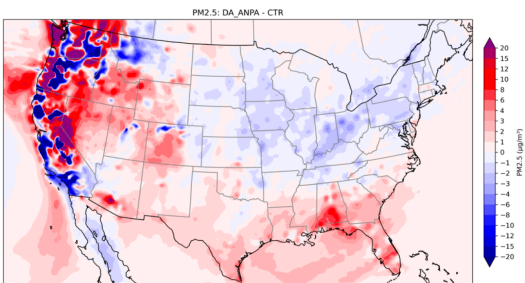


469 (g) DA_ANPA



470

(h) DA_ANPA-CTR



471 **Figure 7.** Spatial distribution of average $PM_{2.5}$ at forecast hour 1, based on 80 forecasts initialized four times daily
 472 (0000 UTC, 0600 UTC, 1200 UTC, and 1800 UTC) during 2–21 September.

473 (a) $PM_{2.5}$ in experiment CTR (shaded) overlaid with AirNow $PM_{2.5}$ observations (filled dots).



474 (b) $\text{PM}_{2.5}$ bias in experiment CTR.

475 (c) $\text{PM}_{2.5}$ in experiment DA_AN (shaded) overlaid with AirNow $\text{PM}_{2.5}$ observations.

476 (d) $\text{PM}_{2.5}$ difference between experiments DA_AN and CTR.

477 (e) $\text{PM}_{2.5}$ in experiment DA_PA (shaded) overlaid with AirNow $\text{PM}_{2.5}$ observations.

478 (f) $\text{PM}_{2.5}$ difference between experiments DA_PA and CTR.

479 (g) $\text{PM}_{2.5}$ in experiment DA_ANPA (shaded) overlaid with AirNow $\text{PM}_{2.5}$ observations.

480 (h) $\text{PM}_{2.5}$ difference between experiments DA_ANPA and CTR.

481

482 Figure 8 shows the percentage change in MAE (%) between the DA experiments and the control (CTR) experiment
483 at forecast hours 1 and 24. Negative values indicate a reduction in MAE. All DA experiments show improvements:
484 at forecast hour 1, MAE is reduced by approximately 60% in DA_AN and DA_ANPA, and by around 18% in
485 DA_PA. At forecast hour 24, reductions are smaller but still present—about 6–7% in DA_ANPA and around 2% in
486 DA_PA. MAE is reduced at nearly all stations at forecast hour 1, and at most stations by forecast hour 24.

487 In summary, DA also improves the spatial distribution of $\text{PM}_{2.5}$ forecasts. Figure 7 shows that the control run
488 exhibits significant regional biases, with overpredictions along the California coast and in parts of the Midwest and
489 Northeast (EPA regions 1, 2, 3, 4, and 7), and underpredictions in the Southwest and Gulf Coast (EPA regions 4 and
490 6). All three DA experiments reduce these errors, particularly correcting coastal overpredictions and improving
491 forecasts in regions affected by wildfires. While DA_PA provides slightly weaker corrections than DA_AN, it
492 contributes meaningful improvements in the Mountain West and Southwest.

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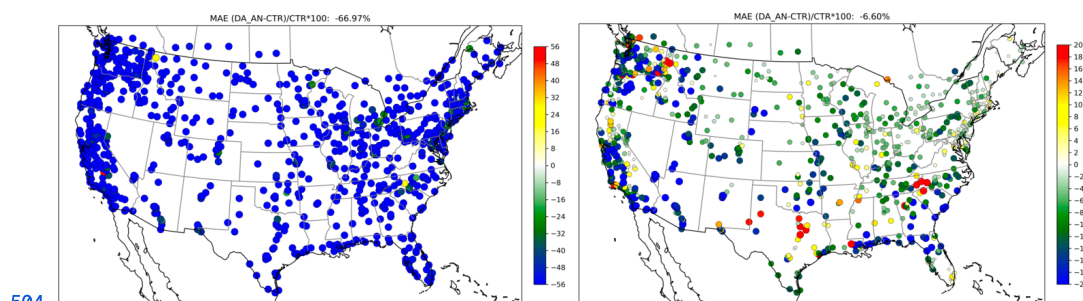
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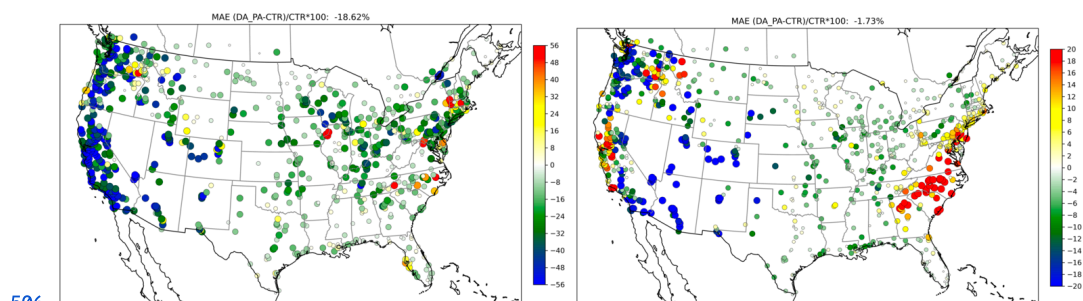


503 DA_AN vs. CTR



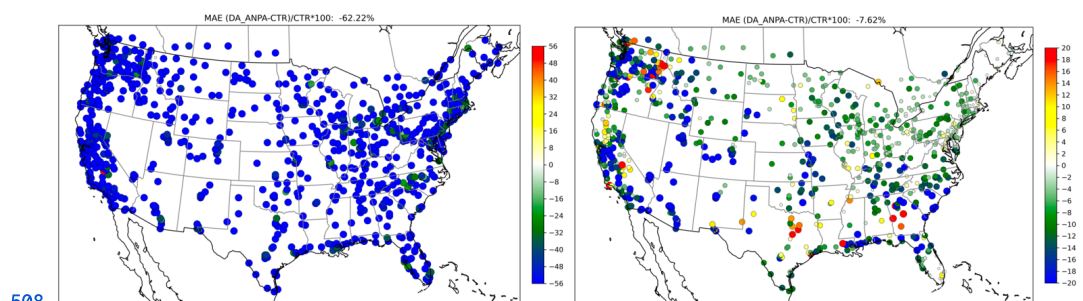
504

505 DA_PA vs. CTR



506

507 DA_ANPA vs. CTR



508

509 **Figure 8.** Percentage change in MAE (%) between DA experiments and the control (CTR) experiment. The
 510 percentage is calculated as $(MAE(DA)-MAE(CTR))/MAE(CTR)*100$.

511 Left panels show the 1-hour forecast; the right panels show the 24-hour forecast.

512 Top row: DA_AN vs. CTR; Middle row: DA_PA vs. CTR; Bottom row: DA_ANPA vs. CTR.

513 4.2. Results from Forecasts Initialized at 1200 UTC

514 In this section, we examine the forecasts initialized at 1200 UTC, which is the time when the operational AQM
 515 launched a 72-hour forecast. Time series of 1–24-hour $PM_{2.5}$ concentrations are analyzed by grouping the EPA
 516 regions into three areas:



- 517 • Area A includes EPA Regions 8, 9, and 10, which are the areas where fires occurred and/or were most
518 influenced by smoke.
- 519 • Area B includes EPA Regions 1, 2, 3, 5, and 7, where the control run generally shows an overall
520 overprediction (see Fig. 6b).
- 521 • Area C includes EPA Regions 4 and 6, which show an overall underprediction in the control run (also
522 shown in Fig. 6b).

523 Time series of 1–24 h $PM_{2.5}$ forecasts for the above three areas are shown in figure 9. There is a large
524 overprediction (spikes) in Area A from September 9 to 11 in the control run, followed by a transition to
525 underpredictions from September 12 to 17. The overprediction is due to inaccurate fire emissions, as similar spikes
526 are not observed in Area B and C. However, this overprediction contributes to spurious “good” performance in the
527 control run during the transition period from September 11 to 12. As shown during this period, the later (~3–24 h)
528 forecasts from the DA_AN and DA_ANPA experiments do not outperform the control run, although their first-hour
529 forecasts are closer to AirNow observations. Overall, the DA experiments clearly improve both the overprediction
530 from September 9 to 11 and the underprediction from September 14 to 17.

531 Regarding the forecasts over Areas B and C, the overprediction in Area B and the underprediction in Area C are
532 generally improved.

533 Figure 10 shows $PM_{2.5}$ forecast error statistics (Bias, MAE, and RMSE) for forecast hours 1–24. At the first
534 forecast hour, the DA experiments in Areas A, B, and C outperform the control run across all three metrics. In terms
535 of MAE and RMSE, DA_AN and DA_ANPA perform better than both DA_PA and the control run in Areas A and
536 B. For the 1–24 h $PM_{2.5}$ forecasts in Area C, all DA experiments outperform the control run in terms of Bias and
537 MAE. However, RMSE improvements are only seen up to forecast hours 7–9.

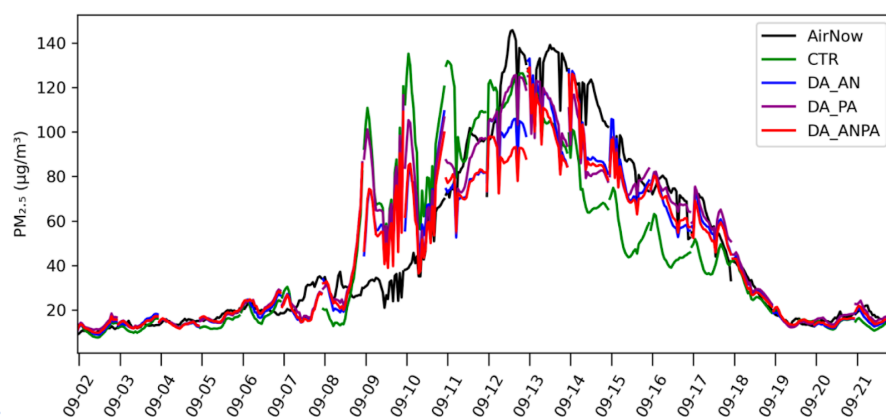
538 The impact of additional assimilation of PurpleAir data shows area-dependent behavior. For example, it slightly
539 reduces MAE and RMSE in Area A and noticeably reduces Bias, MAE, and RMSE in Area B. In Area C, no clear
540 positive impact is observed, although assimilating PurpleAir data alone still results in better performance than the
541 control run.

542 Performance diagrams (Fig. 11) show consistent improvements in Critical Success Index (CSI) scores across all
543 forecast hours for all DA experiments compared to the control run. Among the DA configurations, DA_AN and
544 DA_ANPA show comparable performance and both outperform DA_PA at 1h and 12h forecasts.



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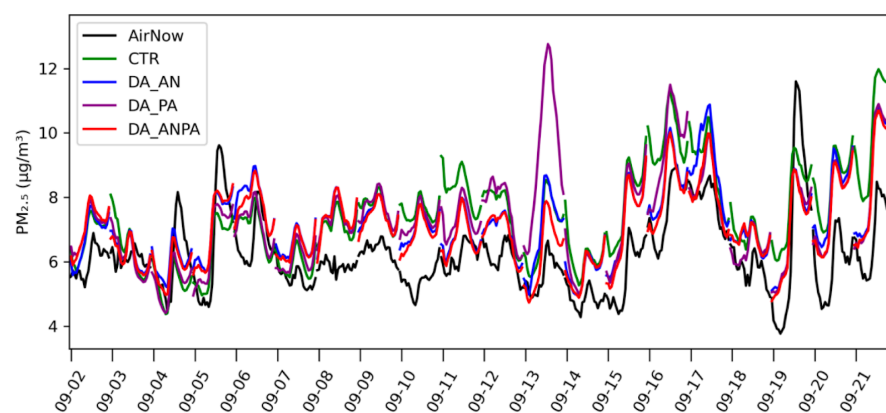
(a)



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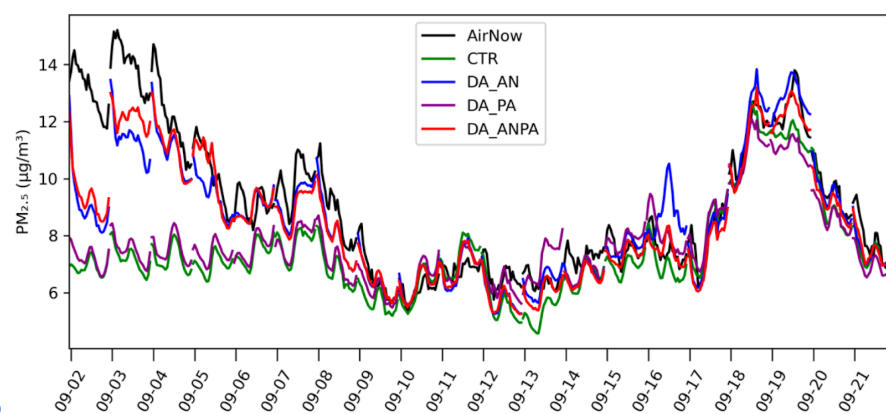
(b)



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(c)

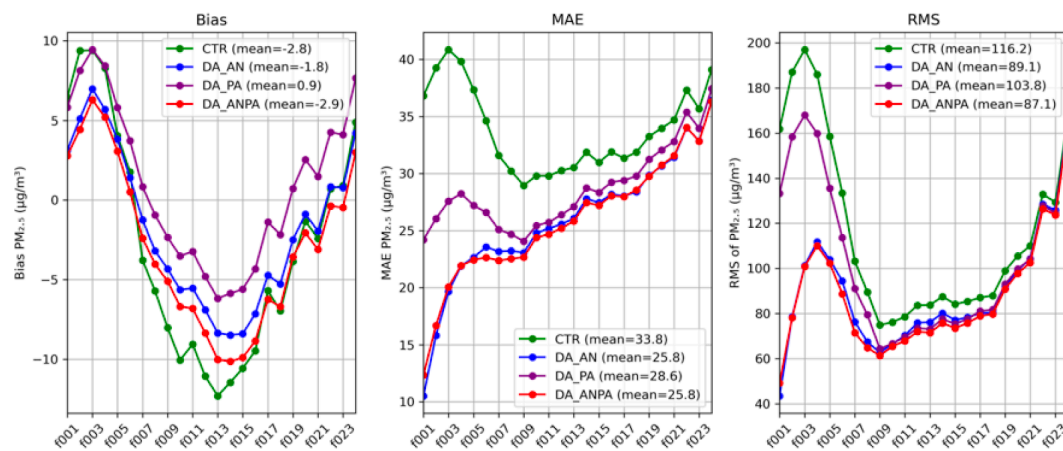


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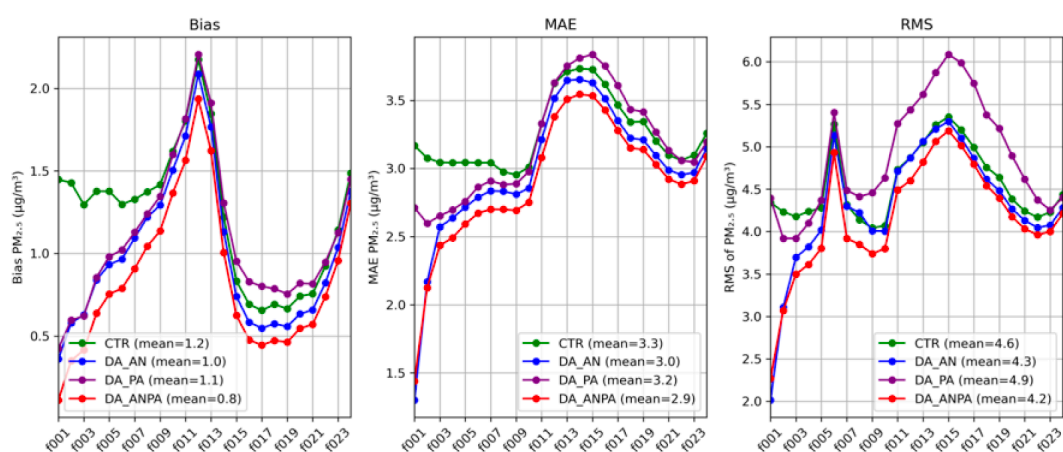
551 **Figure 9.** Time series of 1–24 h $PM_{2.5}$ forecasts: (a) averaged over EPA Regions 8–10; (b) averaged over EPA
 552 Regions 1, 2, 3, 5, and 7; and (c) averaged over EPA Regions 4 and 6.



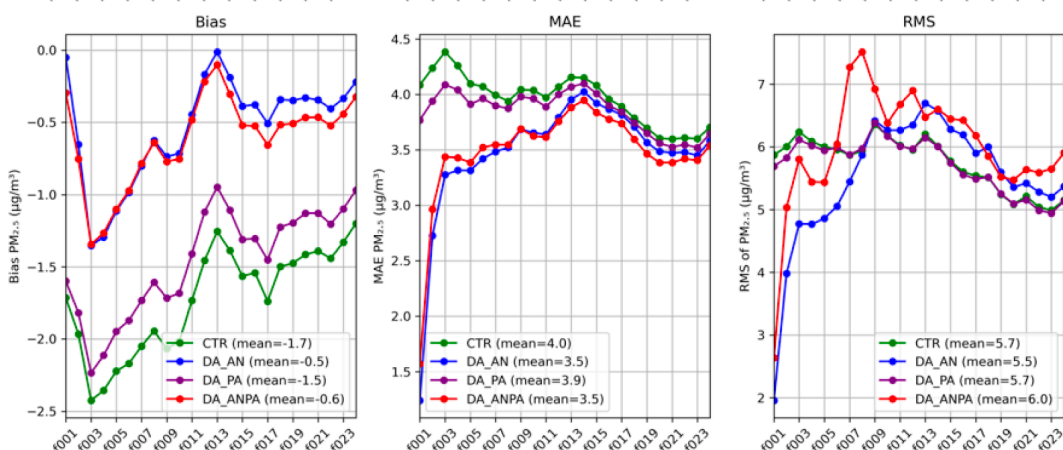
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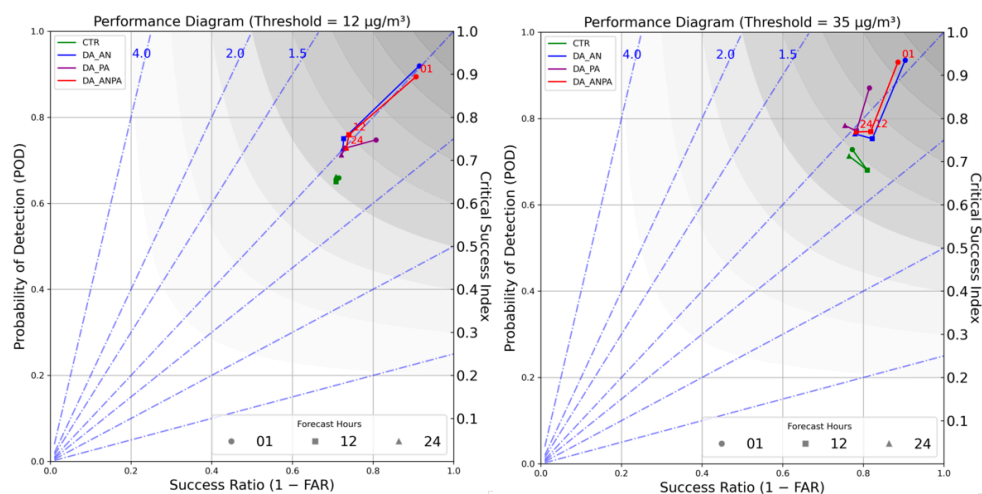


556

557 **Figure 10.** PM_{2.5} forecast statistics at forecast hours 1–24. Top row: averaged over EPA Regions 8–10.; Middle
 558 row: over EPA Regions 1,2,3,5,7, and Bottom row: over EPA Regions 4,6.



559



560

561 **Figure 11.** Performance diagram for forecast hours 1, 12, and 24 with a PM_{2.5} threshold of (a) 12 µg/m³, (b)
 562 35 µg/m³.

563

564 5 Summary and discussion

565 The latest version of NOAA's regional AQM system was implemented and became operational on May 14, 2024.

566 This system has been developed through the online coupling of the Finite Volume Cubed Sphere (FV3) atmospheric
 567 model with the Environmental Protection Agency (EPA)'s Community Multiscale Air Quality (CMAQ) model
 568 within the Unified Forecast System (UFS) framework. In order to provide improved initial conditions for AQM
 569 supporting enhanced prediction of wildfire impacts on air quality prediction, the capability to assimilate PM_{2.5}
 570 observations into AQMv7 was developed within JEDI and tested using its 3D-Var assimilation component. Note that
 571 the developed assimilation scheme can also be used to generate analysis (reanalysis) dataset for other applications,
 572 for example, providing data for training artificial intelligent models used in air quality prediction.

573 Data assimilation experiments were conducted for the September 2020 Western U.S. wildfire episode, using
 574 3-hourly cycling with observations from the AirNow and PurpleAir networks. Results showed that assimilating
 575 AirNow PM_{2.5} observations significantly improved 1–24 hour forecast skill. Mean absolute error (MAE) was
 576 reduced by approximately 60% at forecast hour 1 and 7% at hour 24, relative to forecasts without data assimilation.
 577 Assimilating PurpleAir data alone yielded more modest improvements—approximately 18% at hour 1 and 2% at
 578 hour 24—but when combined with AirNow, PurpleAir data provided additional benefit by further reducing MAE
 579 slightly either compared with AirNow observations (Fig. 3) or PurpleAir observations (Figure not shown). The
 580 positive impact of the PurpleAir data assimilation during the September 2025 wildfires was also demonstrated in an
 581 experimental Rapid Refresh Forecast System coupled with Smoke and Dust Model (Wang et al., 2023), where it
 582 significantly reduced the model's 24-h underprediction of surface PM_{2.5}. Considering the PurpleAir data coverage



583 has improved since September 2025, the results of this study further highlight its potential to complement AirNow
584 observations by filling spatial gaps and improving PM_{2.5} forecast skills.

585 In this first implementation of aerosol data assimilation in JEDI for AQMV7, the control variables are defined as
586 individual forecast aerosol variables. In previous work on aerosol data assimilation for an earlier version of AQM
587 using the GSI system (Wang et al., 2021), one option for the control variables was to define them as the total aerosol
588 mass in each of the three modes, resulting in just three control variables. A control variable transform (CVT) was
589 then applied to partition the analysis increments across these modes to individual aerosol species, based on the ratio
590 of each species' mass to the total mass within the corresponding mode. The use of total aerosol mass in the three
591 modes as control variables—thereby reducing the number of control variables from 70 to 3—is planned for a future
592 phase of development. The use of total masses as control variables also reduces the cost of the background error
593 statistics calculation and iterative minimization. (Kumar et al. 2019). It is noted that the ensemble based data
594 assimilation approach is superior to capture flow-dependent background error covariances and aerosol assimilation
595 along with emission updates can be developed when an ensemble prediction system for AQM is there.

596 This study focused on surface-level PM_{2.5} and did not incorporate vertical profile constraints with satellite-based
597 aerosol optical depth (AOD) retrievals, which could further enhance forecast skill. A key challenge is the need for a
598 robust forward operator in the CRTM AOD module—specifically, the creation and validation of lookup tables
599 (LUTs) for AOD calculations with AQM. As an intermediate solution, existing LUTs in CRTM, such as the GEOS-5
600 LUTs, have been tested by grouping and mapping AQM aerosol species to those used in GEOS-5 (Wang et al.
601 2025). However, this approach presents several issues. For instance, AQM does not distinguish between hydrophilic
602 and hydrophobic aerosol types of organic carbon and black carbon, whereas GEOS-5 does. Additionally, AQM
603 (through CMAQ) uses a modal aerosol representation, while GEOS-5 adopts a bin-based approach, making the
604 mapping between the two systems non-trivial. AOD assimilation also depends on an accurate vertical distribution of
605 aerosols in the background field so that the CRTM AOD operator can provide meaningful gradient information at
606 the correct vertical levels to constrain the analysis update. However, AQM models have shown deficiencies for the
607 September 2020 fire events in representing smoke concentrations at and above plume rise levels, largely due to how
608 fire emissions are injected into the model. This will be improved in the next update of the operational AQM.

609 **Code and data availability**

610 The AQMV7 model, JEDI software and PM_{2.5} and fire emission data we used in this research are publicly available
611 on Zenodo (<https://doi.org/10.5281/zenodo.17049857>; Wang et al., 2025b).

612 Users are referred to the guidance on compiling and running the model:
613 <https://ufs-srweather-app.readthedocs.io/en/develop/UsersGuide/index.html> (Last accessed on August 26, 2025).

614 Global Forecast System analysis data were downloaded from the NCAR Research Data Archive:
615 <https://doi.org/10.5065/D65D8PWK> (last access: Aug 26 2025)



616 Author contribution

617 HW designed and developed the PM_{2.5} DA capability within JEDI for the AQM model, conducted experiments, and
618 evaluated performance; CM and JB contributed to PM_{2.5} DA methodology, advised on code implementation, and
619 assisted in performance analysis; SW contributed to PM_{2.5} DA methodology and experimental design; RL
620 conducted control experiments and contributed to workflow development; JL and KW contribute to model
621 configuration and control run setup; YT contributed to background error modeling and observational error
622 specification; HC, AT and HL contributed to workflow development; JL performed quality control and correction of
623 PurpleAir observations.

624 Competing interests

625 The authors declare that they have no conflict of interest.

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628 the PurpleAir observations.

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631 Cooperative Agreement NA22OAR4320151 with the University of Colorado. The scientific results and conclusions,
632 as well as any views or opinions expressed herein, are those of the authors and do not necessarily reflect those of
633 NOAA or the Department of Commerce.

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