

1 **Residential burning is a potentially significant source of soluble iron to the**
2 **ocean**

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32

33 **Abstract**

34 Understanding the physicochemical processes that supply atmospheric aerosol iron (Fe) to the
35 ocean is crucial for understanding of global biogeochemical cycles. Anthropogenic activity
36 contributes significant fluxes of aerosol Fe to the atmosphere, the soluble fraction of which can
37 modulate marine primary productivity upon its deposition to the ocean surface. However,
38 anthropogenic aerosol Fe solubility remains poorly constrained, due in part to a lack of direct
39 measurements spanning a multitude of anthropogenic sources. We measured solubility of
40 aerosol Fe from several distinct anthropogenic combustion processes and fuel types. The
41 median Fe solubility varied widely by source, ranging from 0.03% for power plant coal fly ash
42 to 55.87% for biofuel burning; furthermore, residential coal burning aerosol possessed much
43 higher Fe solubility than power plant coal fly ash. Using the new Fe solubilities reported herein,
44 we updated parameters for anthropogenic aerosol Fe within the Community Earth System
45 Model. Anthropogenic combustion is estimated to contribute up to 20% of the global soluble
46 Fe flux to the ocean in the present day. Furthermore, we identified residential coal burning as
47 a previously neglected but potentially important source with regional flux contributions
48 ranging from <1% to 21%. Our work underscores the need to further refine understanding of
49 aerosol Fe properties from a wide variety of anthropogenic sources by increasing observations
50 in more novel aerosol regimes, with a focus on residential coal burning. This understanding
51 will in turn aid in characterizing the influences of anthropogenic activity on past, present, and
52 future atmospheric nutrient inputs to marine ecosystems.

53

54 **1 Introduction**

55 Anthropogenic activities have altered the atmospheric burden and deposition fluxes of
56 biogeochemically relevant trace metals, including iron (Fe) (Bergas-Massó et al., 2023;
57 Hamilton et al., 2020b). The quantity of Fe in ocean waters plays a particularly important role
58 in modulating the spatiotemporal distribution of primary productivity in ocean ecosystems,
59 which has downstream impacts on marine fisheries and carbon sequestration (Ito et al., 2021;
60 Tagliabue et al., 2014; Tagliabue et al., 2017). Energy-production, transportation, shipping,
61 and manufacturing (e.g., steel production) are all characterized sources of anthropogenic
62 aerosol Fe (Ito and Miyakawa, 2023; Ito and Shi, 2016; Rathod et al., 2024). These differing
63 combustion fuel types possess distinct physicochemical properties that influence their impact
64 on radiative forcing and nutrient supply (Al-Abadleh et al., 2023; Ito et al., 2018; Matsui et al.,
65 2018; Rathod et al., 2020).

66 To assess the potential nutritional impact of atmospheric Fe deposition on ocean
67 ecosystems, atmospheric aerosol research primarily focuses on tracing the soluble Fe content
68 in aerosol (Baker et al., 2020; Ito et al., 2019; Mahowald et al., 2018). Soluble Fe content is
69 often expressed as the fraction of soluble to total Fe in aerosol and then reported as a percentage
70 solubility (Baldo et al., 2022; Liu et al., 2022; Mahowald et al., 2009). Several key processes
71 control solubility of aerosol Fe over the course of its lifetime: 1) Fe mineralogy, 2) interactions
72 with acidic and organic species in aerosol and cloud water, and 3) particle size and surface area
73 to volume ratios (Bergas-Massó et al., 2023; Journet et al., 2008; Mcdaniel et al., 2019).
74 Anthropogenic combustion not only alters the magnitude and spatial distribution of Fe fluxes
75 from the atmosphere and to the surface ocean, but also influences the composition of the

76 atmosphere, that in turn, influences dissolution chemistry of aerosol Fe both directly and
77 indirectly. Mixing of aerosol Fe with acidic (e.g., sulfates or nitrates) and organic species (e.g.,
78 oxalate) co-emitted during anthropogenic combustion increases its solubility during transport
79 (Bergas-Massó et al., 2023; Chen et al., 2024; Itahashi et al., 2022; Li et al., 2017; Longo et al.,
80 2016). Furthermore, diverse technologies utilized during combustion processes (i.e., variable
81 combustion temperatures, boilers vs. furnaces, degree of emission control, and the fuel quality)
82 also influence the physicochemical properties of aerosol Fe beyond the composition of fuel
83 alone. As a result, studies examining socioeconomic, technology, and policy driven changes to
84 anthropogenic fuel-burning are needed to anticipate impacts on the global Fe cycle (Hamilton
85 et al., 2020b).

86 When compared to mineral dust, anthropogenic emissions of aerosol Fe are several orders
87 of magnitude lower at the global scale; however, anthropogenic Fe has a higher fractional
88 solubility (Ito et al., 2021), and the relative contribution of soluble Fe from anthropogenic
89 combustion is spatially distinct from dust (Hamilton et al., 2020b; Hamilton et al., 2019).
90 Therefore, anthropogenic activity can be a major contributor to Fe fluxes in many high nutrient
91 low chlorophyll (HNLC) ocean regions (Hawco et al., 2025; Liu et al., 2022).

92 Despite the importance of understanding anthropogenic Fe fluxes, the fractional solubility
93 of aerosol Fe emitted from various anthropogenic sources remains poorly understood
94 (Desboeufs et al., 2005; Li et al., 2022b; Oakes et al., 2012); consequently, Fe solubility
95 parameterizations in modeling studies for anthropogenic Fe vary widely (Ito et al., 2019;
96 Myriokefalitakis et al., 2018). In this work, we measured the Fe content and solubility for
97 aerosol emitted by several important anthropogenic sources (i.e., coal power plants, steelwork

98 industry, municipal water combustion, oil combustion, residential coal, and biofuel burning).
99 Then, using an Earth System Model, we applied the experimental results by updating Fe
100 solubility parameters for distinct anthropogenic combustion fuel-sources. Simulated Fe
101 concentrations and solubilities were validated against a global observational Fe aerosol dataset
102 at the regional scale. Then, the model was used to quantify and bound uncertainties in emission
103 and deposition fluxes of soluble Fe under three anthropogenic combustion emission scenarios
104 spanning past (pre-industrial) to future (Shared Socioeconomic Pathway 3-7.0, SSP370)
105 conditions.

106 **2 Methodology**

107 The experimental and modelling methods employed in this work are described in Sections
108 2.1 and 2.2, respectively.

109 **2.1 Experimental methods**

110 This work examined six types of anthropogenic combustion aerosol, which were classified
111 into two broad categories. The first category, fly ash, included power plant coal fly ash,
112 steelwork fly ash, municipal waste fly ash, and oil fly ash. The second category, residential fuel
113 sources, included residential coal and biofuel combustion. Biofuels examined in this work were
114 limited to straw, wood, grasses and leaves, and we did not examine other biofuels such as dung.

115 **2.1.1 Fly ash and bottom ash samples**

116 The volume-mean diameters, determined using diameter light scattering, were found to
117 be 16.9-67.6, 4.7-176.4, 21.2-115.9 and 15.4 μm for power plant coal fly ash (n=31), steelwork
118 fly ash (n=29), municipal waste fly ash (n=3), and oil fly ash (n=1) samples, respectively (Li,
119 2025).

120 Power plant coal fly ash samples were obtained from electrostatic precipitators or
121 baghouse rows in coal power plants in 29 provinces in China (Li et al., 2021; Liu et al., 2021);
122 one coal power plant was selected in each province except for Guangdong and Shandong where
123 two coal power plants were selected for each province. As a result, 31 power plant coal fly ash
124 samples were examined in total. In addition, we examined 29 steelwork fly ash samples
125 collected from different iron and steel plants, three municipal waste fly ash samples (Li et al.,
126 2022b; Li et al., 2021), and two oil fly ash samples which were PM_{2.5} samples emitted by heavy
127 oil and diesel fuel combustion in the engine of a cargo ship (Wu et al., 2018), and one oil
128 bottom ash sample (Fu et al., 2012). As the numbers of municipal waste and oil ash samples
129 were limited, we include their results data in the supplement, but do not discuss them further
130 in the main paper due to a lack of statistical significance.

131 Fly ash and bottom ash samples (~10 mg for each sample) were digested and then
132 analyzed using inductively coupled plasma mass spectrometry (ICP-MS) to determine their Fe
133 content. Experimental procedures for sample digestion and total Fe measurement can be found
134 elsewhere (Li et al., 2022c). Soluble Fe was leached and determined using the procedure
135 described in our previous work (Li et al., 2022b). In brief, fly ash and bottom ash samples (~20
136 mg for each sample) were individually leached in 20 mL sodium acetate buffer (5 mmol/L, pH
137 = 4.3) for 2 h, during which an orbital shaker (300 rpm) was used to stir the solution. The
138 aqueous mixture was centrifuged (3000 rpm) for 15 min, and a pH paper (range: 3.5-6.8;
139 precision: 0.3; Macherey-Nagel, Germany) was used to measure the pH of the solution and no
140 measurable pH change occurred after leaching. The aqueous solution was filtered through a
141 polyethersulfone filter (pore size: 0.22 μm), acidified to contain 1% (v/v) nitric acid and then

142 analyzed by ICP-MS to measure soluble Fe. In this work, fractional solubility of Fe was
143 reported as the ratio (in %) of soluble Fe to total Fe.

144 A wide range of protocols, differing in leaching solution, filter pore size, and so on, were
145 used to extract soluble Fe, and the results obtained using different leaching protocols could be
146 substantial (Tang et al., 2025). Sodium acetate buffer, instead of ultrapure water, was used in
147 the present work as the leaching solution, because its pH did not change during leaching due
148 to its larger buffering capacity compared to ultrapure water.

149 **2.1.2 Residential coal and biofuel combustion aerosols**

150 Generation and collection of residential coal and biofuel combustion aerosols are detailed
151 in the Supplement (Text S1). In brief, we burned coal and biofuel in a commercial cook stove
152 widely used in rural areas in China and collected PM_{2.5} samples (aerosol particles with
153 aerodynamic particle diameters below 2.5 μm) onto pre-cleaned Whatman 41 (W41) cellulose
154 filters using a medium volume aerosol sampler (TH-150C, Tianhong Co.).

155 Our work examined three types of coal (anthracite, semibituminous coal, and bituminous
156 coal) and nine types of biofuel (wheat straw, rice straw, corn straw, rape straw, cogongrass,
157 China fir trunk, pine trunk, poplar trunk, and pine needle) commonly found in China. We
158 collected eight filter samples for each fuel type, except anthracite for which we only collected
159 two filter samples. We had to combine some filter samples in our experimental analysis to meet
160 the detection limit for soluble Fe; as a result, the number of effective filter samples (for which
161 Fe content and solubility were reported) were usually <8 for each fuel type (see Tables S2 and
162 S4 for further information).

163 After aerosol collection, the filters were individually placed in a pre-cleaned Petri dish
164 and then stored in a desiccator for 60 h to remove particle-associated water. The mass of filters
165 before and after aerosol collection were measured (accuracy of 0.1 mg), and the mass of
166 particles collected ranged from 2.5 to 432.7 mg. Each filter was then divided into two equal
167 parts. To determine the soluble Fe content, the first half of a filter was leached in 20 mL sodium
168 acetate buffer (5 mmol/L, pH = 4.3) for 2 h (Section 2.1.1) and analyzed using ICP-MS. Fe
169 concentrations in some leaching solutions were low; as a result, these leaching solutions (~15
170 mL for each solution) were combined for the same fuel type and then pre-concentrated to a
171 volume of 6 mL, in order to increase Fe concentration in the solution used for ICP-MS analysis.
172 The second half of a filter was digested and analyzed by ICP-MS to determine total Fe, and the
173 experimental procedure used can be found in our previous work (Zhang et al., 2022). If leaching
174 solutions were combined for the first parts of these filters, their second parts were also
175 combined and digested together to allow direct comparison.

176 **2.1.3 Quality assurance and quality control**

177 The detection limit of Fe in solution was determined to be 0.5 $\mu\text{g/L}$ in this work. A
178 reference solution (NIST 1643f) was used to check the accuracy of ICP-MS analysis, and the
179 difference between actual and measured concentrations was found to be <1%. Furthermore,
180 three blanks (with no fly ash or filters not loaded with any particles) were used in each batch
181 when we measured total or soluble Fe. The background levels of soluble Fe were always below
182 the detection limit; the background levels of total Fe, ranging from 4.3-5.7 $\mu\text{g/L}$, were much
183 lower than total Fe concentrations for most of our samples and subtracted when we reported
184 our results.

185 **2.2 Model simulations**

186 **2.2.1. Atmospheric Fe model description**

187 Earth System Models can investigate the spatiotemporal distribution and fluxes of key
188 atmospheric nutrients under various climatological regimes (Hamilton et al., 2020b; Hamilton
189 et al., 2022; Wu et al., 2020). To test the impact of new soluble Fe parameters (Section 2.1) on
190 modeled fluxes of soluble aerosol Fe to the atmosphere and marine ecosystems, we used the
191 Mechanism of Intermediate complexity for Modeling Iron (MIMI). MIMI is an Fe aerosol-
192 chemistry module embedded within the atmospheric component (Community Atmosphere
193 Model version 6, CAM6) of the Community Earth System Model version 2 (CESM2)
194 (Danabasoglu et al., 2020; Hamilton et al., 2019). Mineral dust, anthropogenic combustion,
195 and wildfire emissions are currently represented as distinct sources of aerosol Fe in MIMI. The
196 current dust emission scheme within MIMI includes an updated soil moisture submodule
197 within the land component of the model that prognostically calculates dust aerosolization as a
198 function of soil moisture (Li et al., 2022a). Following the implementation of a new soil-
199 moisture scheme, dust was rescaled to attain a global climatological mean dust aerosol optical
200 depth of ~ 0.03 (Ridley et al., 2016), consistent with all previous versions of the MIMI model.
201 The inclusion of these improvements to dust and updated anthropogenic Fe sources represents
202 a new working version of MIMI v1.1, as described herein, and detailed validation efforts are
203 reported in the Supplement (Figure S3 and Table S11).

204 A comprehensive overview of model details and parameters is provided in Hamilton et al.
205 (2019); in brief, MIMI simulates the emission, atmospheric transport, and deposition of Fe-
206 containing aerosol within three distinct particle size modes (Aitken, accumulation, and coarse

207 modes). Within each source of aerosol Fe (dust, wildfire, and anthropogenic combustion), both
208 the insoluble and soluble fractions are carried as separate tracers, and the soluble fraction of Fe
209 for each aerosol source is assigned at the point of emission. Prior to deposition and during
210 atmospheric transport, Fe solubility is further modified via non-reversible multiphase reactions
211 with acidic and organic species. Acidic processing is a function of aerosol pH and temperature
212 (Meskhidze et al., 2005), while organic processing is an aqueous phase chemistry reaction that
213 depends on oxalate concentrations which are calculated based on the concentrations of
214 secondary organic aerosol present (Johnson and Meskhidze, 2013; Scanza et al., 2018).

215 The model is gridded in a 3-dimensional space at a resolution of 0.96×1.25 degrees
216 (latitude \times longitude) and includes 56 vertical pressure levels from the surface to 2 hPa at the
217 highest altitude. Meteorology is forced in all the simulations using Modern-Era Retrospective
218 analysis for Research and Applications Version 2 (MERRA-2), and a 1-year model spin up
219 was undertaken for all simulations.

220 **2.2.2 Global pyrogenic Fe emission inventories and input dataset development**

221 While dust Fe emissions are calculated prognostically within MIMI, anthropogenic and
222 wildfire (sum of these being pyrogenic) emissions are prescribed using emissions inventories.
223 Annual mean anthropogenic Fe emission fluxes were inputted to the model using a modified
224 version of an inventory first developed in Rathod et al. (2020) and further detailed in Rathod
225 et al. (2024). In this inventory, Fe content in combustion aerosol was empirically derived for
226 the present day (PD; climatological year 2010) using the Speciated Pollution Emissions Wizard
227 (SPEW) (Bond et al., 2007; Bond et al., 2004), which characterizes anthropogenic emissions
228 of particulate matter by fuel-source and combustion technology. Soluble and insoluble Fe

229 content are dependent on fuel-type and also segregated by key sectors: 1) industrial fossil fuel
230 (coal), 2) industrial and vehicular fossil fuels (oil), 3) smelting operations (steel/iron), 4)
231 cooking/heating using biomass (biofuel/wood), and 5) waste burning (Rathod et al., 2020).
232 Industrial oil emissions were separated by land- and sea-based emissions to distinguish
233 terrestrial transportation from shipping. Wildfire-Fe emission parameters are detailed in
234 Hamilton et al. (2019), and in this work we used the CMIP6 (Coupled Model Intercomparison
235 Project phase six) fire emission inventory for PD simulations (Van Marle et al., 2017).

236 Here, for the first time, we separated anthropogenic coal Fe into distinct industrial and
237 residential sources and tested three approaches to constrain the magnitude of the residential Fe
238 signal, i.e. a high, central-high, and central-residential emission scenario, with the naming
239 conventions relative to the residential Fe emissions reported by Rathod et al. (2020, 2024).
240 Each inventory increases in its respective source grouping complexity to evaluate key
241 uncertainties in the representation of Fe emissions from residential combustion in global
242 inventories (Figures 1, S4a and S4b; Table S8). Key to this process is the mapping fuel-specific
243 Fe emissions to CMIP6 anthropogenic sectors, enabling future projections of sector-specific
244 emission changes under the SSP scenarios. The CMIP6 anthropogenic emission dataset as
245 detailed in Hoesly et al. (2018) classifies anthropogenic emission sectors as: 0) agriculture, 1)
246 energy, 2) industrial, 3) terrestrial transportation, 4) residential/commercial/other, 5) solvents
247 production/application, 6) waste, and 7) international shipping.

248 To create the first emission inventory ('high-residential'), we multiplied a series of
249 spatially-resolved ($1^\circ \times 1^\circ$ grid box) ratios of residential-to-industrial black carbon (BC)
250 emissions to the Rathod et al. (2024) Fe inventory emissions for coal. Hoesly et al. (2018) do

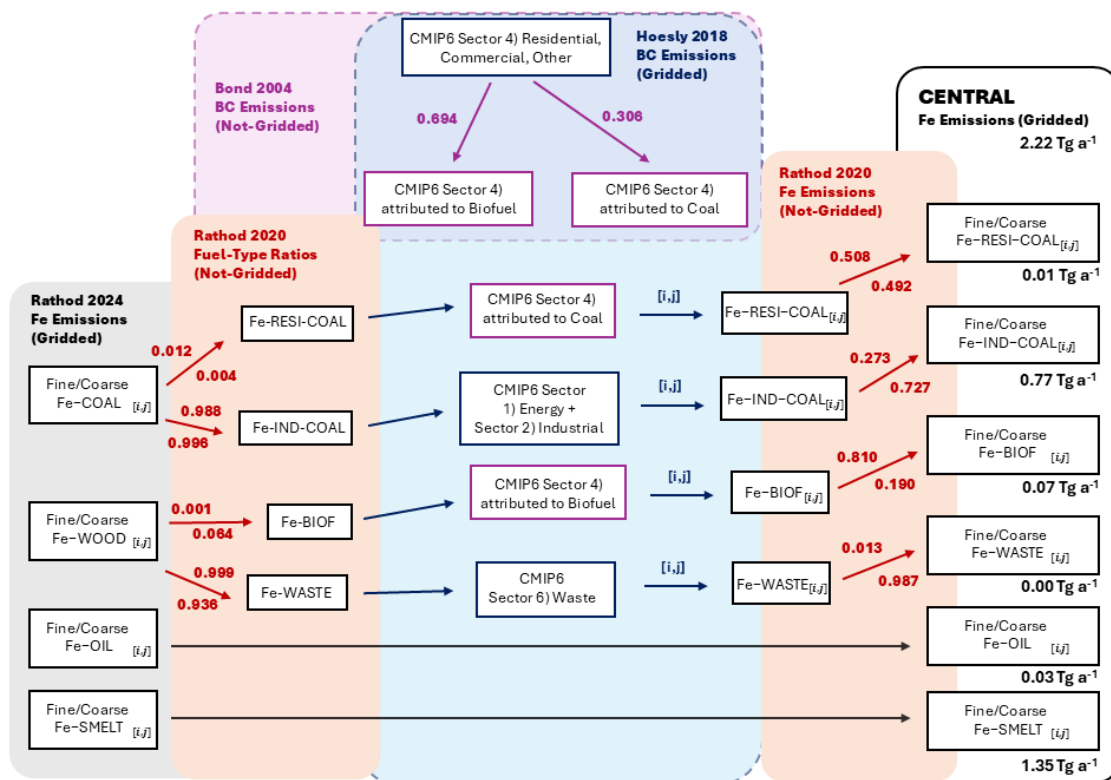
251 not separate biofuel (wood) from residential coal (Table S8); in order to reflect an upper limit
 252 for residential coal burning, we treated all of CMIP6 Sector #4 (residential, commercial, and
 253 other) BC as the residential coal fraction and CMIP6 Sectors #1+2 (energy + industrial) BC as
 254 the industrial coal fraction (Figure S4a; Eqs. 1a and 1b) (Hoesly et al., 2018):

$$255 \quad [Fe_{RESI}]_{i,j,HIGH} = \frac{[BC_{RESI}]_{i,j,H}}{[BC_{RESI+IND}]_{i,j,H}} \times [Fe_{IND+RESI}]_{i,j,R} \quad (1a)$$

$$256 \quad [Fe_{IND}]_{i,j,HIGH} = \frac{[BC_{IND}]_{i,j,H}}{[BC_{RESI+IND}]_{i,j,H}} \times [Fe_{IND+RESI}]_{i,j,R} \quad (1b)$$

257 Where i and j represent the longitudinal and latitudinal coordinates, $RESI$ and IND represent
 258 residential and industrial sources, $HIGH$ represents the high-residential inventory being
 259 constructed, H represents the Hoesly et al. (2018) dataset, R represents the Rathod et al. (2024)
 260 dataset, and $[Fe]$ and $[BC]$ represent their respective speciated fluxes in units of $\text{kg m}^{-2} \text{s}^{-1}$.
 261 This approach was chosen to capture spatial variations in coal burning technologies within the
 262 inventory but assumed that residential Fe-to-industrial Fe emissions track residential BC-to-
 263 industrial BC by grid cell.

264 To create the second emission inventory (central-high-residential') we globally scaled the
 265 high-residential inventory to reflect global Fe emission budgets from residential coal burning
 266 sources previously reported in Rathod et al. (2020). Using the ratio of residential coal to all
 267 residential sources reported in Bond et al. (2004) and assuming parity between BC datasets, we
 268 estimated the portion of CMIP6 Sector #4 BC that could be attributed to residential coal
 269 burning (Table S8). Then, we redistributed the Fe emissions in the high inventory between
 270 residential and industrial coal to reflect this using a global scaling factor of 0.035
 271 (residential:industrial), constituting the central-high inventory (Figure S4b).



272
 273 **Figure 1.** A flowchart representing the steps followed and datasets leveraged to create the
 274 central- anthropogenic Fe emissions inventory. Dashed lines indicate a BC dataset, and solid/no
 275 outlines indicate an Fe dataset. Two values are provided with each red line to show the
 276 fractional split between fine and coarse fractions; the above arrow indicates the fine fraction,
 277 while the below values indicates the coarse fraction. Similar flowcharts for the central-high
 278 and high- inventories are provided in Figures S4a and S4b.

279
 280 To create the third emission inventory ('central-'), rather than using the Hoesly et al. (2018)
 281 BC data to separate residential from industrial coal emissions, we determined the fractional
 282 contribution of residential + commercial coal-Fe to total coal-Fe emissions reported in Rathod
 283 et al. (2020) (commercial and residential sources were specifically grouped to track CMIP6
 284 Sector #4; Figure 1). This separation was applied uniformly at the global scale for each of the

285 fine and coarse aerosol modes individually. The sector-specific ratio of Fe:BC was then applied
 286 to map spatial heterogeneity in emissions following the CMIP6 emissions data (Figure 1),
 287 where sector-specific BC was taken from Hoesly et al. (2018). Sector #4 (Residential,
 288 Commercial, Other) BC emissions were again split into residential coal vs. biofuel sources
 289 using Bond et al. (2004) (Figure 1; Table S8). To best track CMIP6 groupings, we treated
 290 ‘energy’ and ‘industrial’ sources together as industrial coal, ‘waste’ sources as waste, and
 291 ocean-masked ‘international shipping’ with land-masked ‘terrestrial transportation’ together as
 292 oil BC sources. Finally, fine and coarse mode Fe emissions were re-distinguished following
 293 the fuel-specific global averages reported in Rathod et al. (2020) (Figure 1).

294 The central-, central-high, and high-residential coal emissions span three orders of
 295 magnitude at the global scale (4.9, 16, and 460 Gg a⁻¹, respectively; Table S9). Each inventory
 296 was applied and tested with each PD model simulation to perform a sensitivity analysis that
 297 quantified the uncertainty in residential coal Fe emissions introduced by each new inventory.
 298 To capture the lower and upper bounds of uncertainty associated with residential-Fe emissions,
 299 herein we present results from the high and central- inventories only; results from the central-
 300 high inventory are provided in the supplement.

301 Once sector-specific emissions inventories were read into the model, fuel-sources were
 302 summed to total one anthropogenic tracer. This tracer is then separated into soluble and
 303 insoluble fractions with three modes within the model code (Eqs. 2a and 2b). This results in six
 304 anthropogenic combustion-Fe tracers in total to be transported within the model, as follows:

$$305 \quad [Fe_{insol}]_a = \Sigma\{[Fe_X]_{i,j,a,b} * (1 - sol_b)\} \quad (2a)$$

$$306 \quad [Fe_{sol}]_a = \Sigma ([Fe_X]_{i,j,a,b} * sol_b) \quad (2b)$$

307 where X denotes the emissions scenario, a represents the aerosol mode (fine or coarse), b
 308 represents the fuel-source (industrial oil, industrial coal, residential coal, biofuel, or smelting),
 309 $insol$ represents the insoluble fraction, sol represents the soluble fraction, and sol_b represents
 310 the fractional solubility for each fuel-source (b). As a final step, the fine mode was split into
 311 accumulation and Aitken modes by applying a ratio of 9:1.

312 **2.2.3 Model simulations performed**

313 Thirteen model simulations were performed to evaluate the impact of anthropogenic
 314 combustion aerosol solubility updates (Section 2.1) on atmospheric soluble Fe fluxes to key
 315 marine ecosystems (Table 1). For all simulations, we set the model climatology to present-day
 316 (PD) conditions, spanning 2009-2011 inclusive. Simulations were distinguished as cases
 317 (variable Fe solubility parameterizations) within different emission scenarios (variable
 318 anthropogenic combustion emission fluxes).

319

320 **Table 1.** Description of model simulations performed using MIMI with emission scenarios and
 321 emission inventories either directly input (Fe) to the model or utilized to generate the Fe
 322 inventory (BC). PD = present day (2010 CE), PI = pre-industrial (1750 CE), SSP370 = Shared
 323 Socioeconomic Pathway scenario 3-7.0, MID = midcentury (2040-2050 CE) and END = end
 324 century (2090-2100 CE). NA = assumed industrial activity is zero at 1750 CE. Note the BC
 325 budget varies for the central- inventory per the inclusion of waste sources.

Emissions Scenario	Simulation	Fe Emissions Inventory	BC Emissions database	BC Emission (Tg a ⁻¹)
PD	PD-BASE	Rathod et al., 2024	Hoesly et al., 2018	6.46
PD	PD-RESI	High-Residential	Hoesly et al., 2018	6.46
PD	PD-BIOF	High-Residential	Hoesly et al., 2018	6.46

PD	PD-IND	High-Residential	Hoesly et al., 2018	6.46
PD	PD-BASE	Rathod et al., 2024	Bond et al., 2004 & Hoesly et al., 2018	6.46
PD	PD-RESI	Central-Residential	Bond et al., 2004 & Hoesly et al., 2018	6.97
PD	PD-BIOF	Central-Residential	Bond et al., 2004 & Hoesly et al., 2018	6.97
PI	PI-BASE	Hamilton et al., 2020a	NA	NA
PI	PI-BIOF	Hamilton et al., 2020a	NA	NA
FU (2050)	MID-SSP370-BASE	High-Residential	SSP3.70	8.30
FU (2050)	MID-SSP370-BIOF	High-Residential	SSP3.70	8.30
FU (2100)	END-SSP370-BASE	High-Residential	SSP3.70	6.33
FU (2100)	END-SSP370-BIOF	High-Residential	SSP3.70	6.33

326

327 The first seven simulations (i.e., PD simulations) aimed to assess the impact of each new
328 solubility parameter on the ability of the model to capture ship-based observations of total Fe,
329 soluble Fe, and Fe solubility. These simulations applied the high-residential emissions
330 inventory (Section 2.2.2) and were repeated using the central-residential Fe inventory with the
331 exception of PD-IND (Table 1). The first PD case in each set (i.e., PD-BASE) served as a
332 baseline, i.e., no changes were made to solubility when compared to previous studies using
333 MIMI. Residential coal was distinguished from industrial coal emissions, but this had no
334 impact on soluble Fe fluxes since the same fractional solubility (0.2%) was applied to both
335 sources (Rathod et al., 2020; Rathod et al., 2024). In the next three PD cases (PD-RESI, PD-
336 BIOF, PD-IND), fractional solubility was updated incrementally for individual sectors to
337 assess fuel-type specific impacts to soluble Fe fluxes, which are later detailed in Section 3.3;
338 results from PD-IND are reported in the Supplement. Information on model validation and
339 constraint to ship-based observations of aerosol Fe is provided in Section 2.2.5.

340 Using both pre-industrial (PI; 1750 CE) and future (FU; 2050 and 2100 CE) anthropogenic
341 emissions scenarios, we performed six model simulations to examine the impacts of changes
342 in anthropogenic activity on Fe fluxes over time. In each pair of simulations, we applied
343 the -BASE solubility parameters to establish a baseline and the -BIOF solubility parameters to
344 examine an upper bound for residential soluble Fe fluxes. Accordingly, we also used the high-
345 residential Fe emissions inventory framework in development of the FU emissions inventory
346 to further establish a maximum estimate for anthropogenic soluble Fe through the end of the
347 21st century. To isolate how changes in soluble aerosol Fe fluxes responded to changes in
348 emission parameterizations and subsequent dissolution chemistry in the model, PI and FU
349 simulations were conducted with meteorological and climatological conditions identical to the
350 PD (2009-2011).

351 The PI simulations (PI-BASE and PI-BIOF) served as a reference point for comparison to
352 PD and FU simulations, per minimal influence on the Fe cycle by anthropogenic emissions
353 (Table 1). MID-SSP370-BASE, MID-SSP370-BIOF, END-SSP370-BASE, and END-
354 SSP370-BIOF were conducted to evaluate the projected impact of population increases and
355 socioeconomic changes to energy production and fuel-usage over the course of the century.
356 We selected the highest air pollution emissions scenario of the Shared Socio-economic
357 Pathways (SSPs) as detailed in SSP 3-7.0 (SSP370), the “regional rivalry” scenario, which
358 represents anticipated sociopolitical and environmental changes resulting in an increase to
359 radiative forcing by 3-7.0 W m⁻² by the year of 2100 (Riahi et al., 2017). Given that BC
360 emissions are anticipated to peak in the midcentury (2040-2050) but return to PD-comparable
361 emissions by 2100 (Turnock et al., 2020), we assessed projected changes to Fe emissions at

362 both the mid-point (2050) and endpoint of the 21st century (2100). Dust fluxes in future Fe
363 emission scenarios were adjusted to account for dust-climate feedback using a scaling factor
364 ranging between 1.0-1.1, as described in Hamilton et al. (2020a).

365 **2.2.4 Preindustrial (PI) and future (FU) Fe emission estimates**

366 For PI simulations, we used a pre-developed Fe combustion emission inventory (Hamilton
367 et al., 2020b). Only residential (wood) biofuel burning served as an anthropogenic source of
368 Fe due to a presumable lack of industrialized anthropogenic emissions (i.e., fossil fuels and
369 smelting; (Hamilton et al., 2020a); global emission for anthropogenic combustion was 0.7×10^3
370 Gg Fe a^{-1} and only occupied the fine aerosol mode (i.e., sum of Aitken and accumulation
371 modes). Details on the development of the PI Fe combustion emission inventory are provided
372 in Hamilton et al. (2020b).

373 For FU simulations, we developed two new Fe emissions datasets (for 2050 and 2100)
374 which were both derived from our high-residential emissions inventory developed for the PD
375 simulations. Fe emissions were linearly scaled for all combustion sources according to
376 projected changes in anthropogenic BC emissions via the decadal CMIP6 anthropogenic BC
377 emission dataset for 2040-2050 and 2090-2100 (Hoesly et al., 2018; Riahi et al., 2017). BC
378 emissions labeled ‘residential, commercial and other’ were separated into residential coal and
379 residential biofuel sources of BC based on the grid-cell specific ratios of residential coal Fe to
380 residential biofuel Fe in our high-residential emissions inventory.

381 Following a similar approach to the PD high-residential emissions inventory, using a grid-
382 cell resolved dynamic ratio of Fe-to-BC, that was grouped based on sector and aerosol size

383 fraction, we calculated FU Fe emissions tracking CMIP6 projected BC emissions (Eq. 4), as
384 follows:

$$385 \quad \frac{[Fe_X]_{i,j,a,b}}{[BC_X]_{i,j,a,b}} = \frac{[Fe_{PD}]_{i,j,a,b}}{[BC_{PD}]_{i,j,a,b}} \quad (4)$$

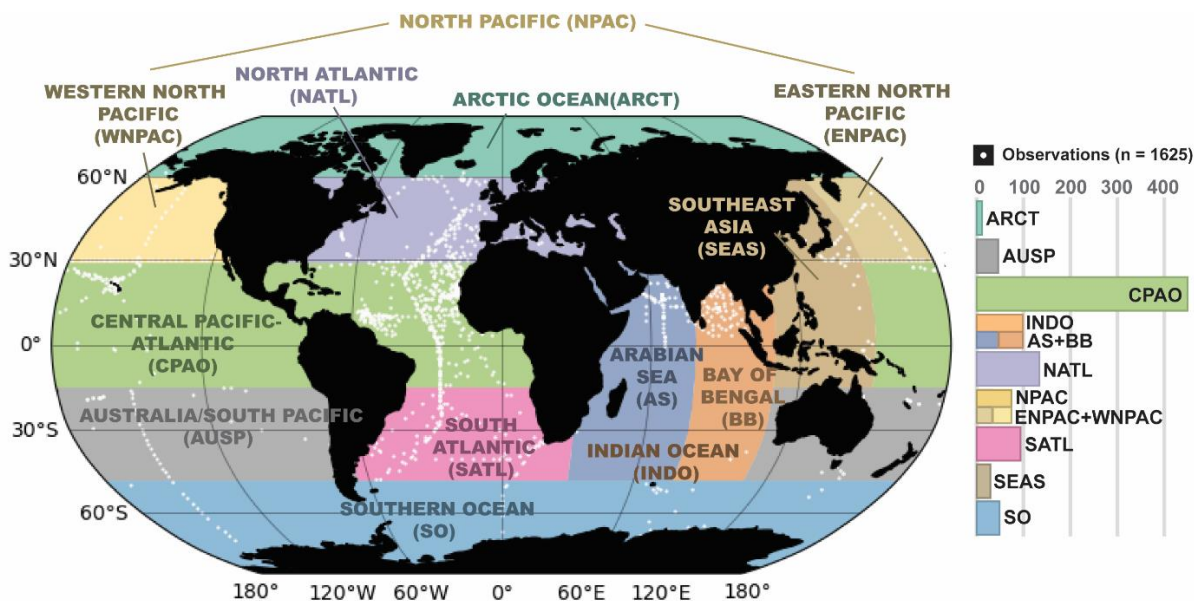
386 where X denotes the emissions scenario (MID-SSP370 or END-SSP370), *i* and *j* represent the
387 longitudinal and latitudinal coordinates, *a* represents the aerosol mode (fine or coarse), *b*
388 represents the fuel-source (industrial oil, industrial coal, residential coal, biofuel, or smelting),
389 and [Fe] and [BC] represent the speciated fluxes ($\text{kg m}^{-2} \text{s}^{-1}$). BC emissions from smelting
390 operations were not directly available for PI or FU projections; therefore, they were set to 0.0
391 in the PI and maintained at PD levels in the FU. By using SSP370, the -BIOF case solubility
392 parameters, and the high-residential emissions inventory, our FU simulations established an
393 upper bound estimate for future anthropogenic soluble Fe fluxes as based on current
394 observational uncertainties. The SimFire inventory, coupled to the LPJ-GUESS (Lund-
395 Potsdam-Jena General Ecosystem Simulator) vegetation model, was used to prescribe wildfire-
396 Fe emissions during the PI era (Hamilton et al., 2018; Hamilton et al., 2020a; Knorr et al., 2016)
397 For wildfire-Fe emissions in FU scenarios, we used the CMIP6 fire emission datasets for MID-
398 SSP370 and END-SSP370 (Bergas-Masso et al., 2025; Hamilton et al., 2024).

399 **2.2.5 Model validation**

400 To evaluate model performance, we compared global observations of total Fe
401 concentration, soluble Fe concentration, and Fe solubility to modeled values for each PD
402 simulation, grouping data by key aerosol deposition and ocean biogeochemistry regions. The
403 observational dataset of Fe content in aerosol was reported in Hamilton et al. (2019) and
404 updated herein to include measurements from Srinivas et al. (2012) and more recent studies

405 published between 2021 and 2024 (n = 1624) (Desboeufs et al., 2024; Elliott et al., 2024;
406 Kurisu et al., 2021; López-García et al., 2021; Marafante et al., 2024; Panda et al., 2022; Perron
407 et al., 2022; Rodríguez et al., 2021; Sakata et al., 2022; Seo and Kim, 2023; Winton et al., 2022;
408 Wu et al., 2023; Zhang et al., 2024).

409 Observed Fe solubility in aerosol spans five orders of magnitude (Perron et al., 2024), and
410 one reason for this large range is due to differences in experimental procedures during
411 quantification (Tang et al., 2025). To facilitate a more direct comparison between modeled and
412 observed soluble Fe content, we removed observations from the global dataset that did not
413 measure soluble Fe directly. When multiple observations fell within a model grid cell, values
414 were aggregated to climatological averages, using medians to be most representative of
415 expected variations in Fe fluxes across time and space (final n = 990; Figure 2). For final
416 evaluation of the model capability in simulating surface Fe concentrations, both model and
417 observational data were grouped into key ocean regions (Figure 2), based on predominant
418 sources of atmospheric aerosol and phytoplankton nutrient limitation dynamics (i.e., HNLC
419 regions) as revealed in Hamilton et al. (2019) and Hamilton et al. (2023). To quantitatively
420 evaluate model skill, root mean square errors (RMSE) were calculated for the high-residential
421 inventory cases and are provided in the Supplement.



422

423 **Figure 2.** Regional groupings for model-observation comparisons of surface Fe concentrations
 424 (ship-based, in aerosol). The coordinates for individual Fe observations are indicated with a
 425 white circle. Number of soluble Fe observations within each region are provided by the
 426 histogram (n=990).

427

428 **3 Results and Discussion**

429 Sections 3.1 and 3.2 present Fe content and solubility measured in our experimental work.
 430 The numbers of samples examined in our work are very small for municipal waste fly ash and
 431 oil fly/bottom ash, and the results may not be representative; therefore, these results are mainly
 432 presented in the Supplement. Modeling results are presented in Section 3.3.

433 **3.1 Fe content by fuel type**

434 This work quantified the Fe content in particles from six different combustion and
 435 anthropogenic sources, including power plant coal fly ash, residential coal combustion aerosol,
 436 steelwork fly ash, residential biofuel burning aerosol, municipal waste fly ash, and oil fly ash
 437 (Table 2; Fe content in individual samples is provided in Tables S1-S5).

438

439 **Table 2.** Summary of Fe content and solubility for power plant coal fly ash, residential coal
 440 combustion aerosol, steelwork fly ash, and biofuel burning aerosol examined in our work (*n*:
 441 number of samples examined in our work). Fe content and solubility for other samples
 442 examined in this work (i.e. municipal waste fly ash, oil fly ash and oil bottom ash) can be found
 443 in the Supplement.

sample type	<i>n</i>	range	average	median
Fe content (mg/g)				
power plant coal fly ash	31	20.7-103.8	37.2±16.8	35.0
residential coal combustion aerosol	10	0.025-0.101	0.044±0.023	0.038
steelwork fly ash	29	5.8-918.9	312.6±246.1	346.5
biofuel burning aerosol	27	0.002-0.101	0.023±0.026	0.013
Fe solubility (%)				
power plant coal fly ash	31	0.002-0.17	0.05±0.05	0.03
residential coal combustion aerosol	10	7.03-100	33.30±27.71	28.45
steelwork fly ash	29	0.007-10.64	1.37±2.77	0.07
biofuel burning aerosol	28	2.86-100	56.07±30.95	55.87

444

445

446 3.1.1 Power plant coal fly ash

447 Fe content ranged from 20.7 to 103.8 mg/g for the 31 power plant coal fly ash samples
 448 examined in our work, with average and median values being 37.2 ± 16.8 and 35.0 mg/g,
 449 respectively. As shown in Table S6, Fe content ranged from 16.0 to 52.0 mg/g (*n* = 3) in one

450 study (Baldo et al., 2022), with mean and median values being 33.0 ± 18.0 and 31.0 mg/g; in
451 another study (Goodarzi, 2006), the median value of Fe content was determined to be 34.4
452 mg/g ($n = 7$). Fe content measured by these two studies (Baldo et al., 2022; Goodarzi, 2006)
453 agreed well with our work. Some other studies (Dutta et al., 2009; Fu et al., 2012; Jankowski
454 et al., 2006; Meij, 1994) found higher mean or median Fe content for power plant coal fly ash
455 (Table S6), but the reported ranges overlapped with our work. For example, Fe content were
456 found to range from 38.3 to 98.6 mg/g ($n = 7$) in one study (Li et al., 2022b), with mean and
457 median values being 62.1 ± 26.7 and 43.2 mg/g; in another study (Moreno et al., 2005), Fe
458 content were found to range from 18.2 to 112.0 mg/g ($n = 23$), with mean and median values
459 being 57.8 ± 22.7 and 52.5 mg/g.

460 In summary, the mean or median Fe content reported in different studies are typically in
461 the range of 30 - 70 mg/g for power plant coal fly ash, and this variability is likely due to
462 difference in coal (Wang et al., 2015; Ward, 2016) and combustion conditions (Blissett and
463 Rowson, 2012; Kutchko and Kim, 2006). Fe content in power plant coal fly ash was set to ~ 70
464 mg/g in some modeling studies (Luo et al., 2008; Rathod et al., 2020), being consistent with
465 experimental results.

466 **3.1.2 Residential coal combustion aerosol**

467 For the 10 residential coal combustion aerosol samples ($PM_{2.5}$) we examined, Fe content
468 ranged from 0.025 to 0.101 mg/g (Table 2), with average and median values being $0.044 \pm$
469 0.023 and 0.038 mg/g, respectively. Only a few previous studies measured Fe content in
470 residential coal combustion aerosols (Table S6). The average Fe content was determined by
471 Patil et al. (2013) to be 0.048 ± 0.035 mg/g ($n = 3$) for $PM_{2.5}$ and 0.061 ± 0.044 mg/g ($n = 3$)

472 for PM₁₀, being similar to or slightly higher than our result. In another two studies (Watson et
473 al., 2001; Zhang et al., 2012), the average Fe content was measured to be 0.671 ± 0.023 mg/g
474 ($n = 4$) and 0.7 ± 0.1 mg/g ($n = 5$), significantly higher than our result, and such differences
475 may be attributed to variations in coal types and combustion conditions. Overall, our and
476 previous studies suggest that the Fe content in residential coal combustion aerosols is very low,
477 typically below 1 mg/g. Fe content were set to 1 and 0.5 mg in previous modeling studies (Luo
478 et al., 2008; Rathod et al., 2020), being broadly consistent with experimental results.

479 Fe content in power plant coal fly ash is much higher than residential coal combustion
480 aerosols, primarily due to differences in combustion conditions (Rathod et al., 2020). Power
481 plant coal fly ash has very low carbon content and is mainly composed of metals and minerals
482 (Ahmaruzzaman, 2010; Li et al., 2022c; Patil et al., 2013); in contrast, residential coal
483 combustion aerosol particles contain a large fraction of carbonaceous materials due to
484 incomplete combustion, and thus the content of metals, including Fe, are much lower (Patil et
485 al., 2013; Zhang et al., 2012). Furthermore, combustion temperature typically ranges from 1200
486 to 1700 °C for coal-fired power plant, enabling Fe in coal to enter fly ash particles through
487 volatilization-condensation (Blissett and Rowson, 2012); residential coal combustion occurs at
488 much lower temperatures which are insufficient for Fe to enter aerosols through this process
489 (Rathod et al., 2020), also leading to lower Fe content.

490 **3.1.3 Steelwork fly ash**

491 For the 29 steelwork fly ash samples we examined, Fe content ranged from 5.8 to 918.9
492 mg/g, with mean and median values measured to be 312.6 ± 246.1 and 346.5 mg/g, respectively
493 (Table 2). As shown in Table S6, some previous studies have reported average Fe content to

494 be 358.9 (n = 1), 369.3 (n = 1), 312.2 (n = 1), and 329.1 ± 22.6 mg/g (n = 4) (Alizadeh and
495 Momeni, 2016; Silva et al., 2019; Souza et al., 2010; Vieira et al., 2013), in good agreement
496 with our results. Lower Fe content was also reported by previous work, with average values
497 being 86.0 (n = 1), 128.1 (n = 1), 150.8 (n = 1), 286.5 (n = 1), 284.6 (n = 1), 238.7 (n = 1), and
498 267.3 ± 4.8 mg/g (n = 4) (Al-Negheimish et al., 2021; Alsheyab and Khedaywi, 2016; Laforest
499 and Duchesne, 2006; Li et al., 2023; Loaiza et al., 2017; Stathopoulos et al., 2013; Xia and
500 Picklesi, 2000); in contrast, some previous studies also found the average or mean Fe content
501 to be around 400-500 mg/g (Machado et al., 2006; Patil et al., 2013; Ye et al., 2021), slightly
502 higher than our results.

503 Despite some variability in Fe content reported by our and previous studies (Table S6),
504 the mean or median Fe content are generally around 300–500 mg/g for steelwork fly ash. In a
505 recent modeling study (Rathod et al., 2020), the Fe content in steelwork fly ash was set to 440
506 mg/g (and the lower and upper bounds were set to 150 and 950 mg/g), being consistent with
507 experimental results.

508 **3.1.4 Biofuel burning aerosol**

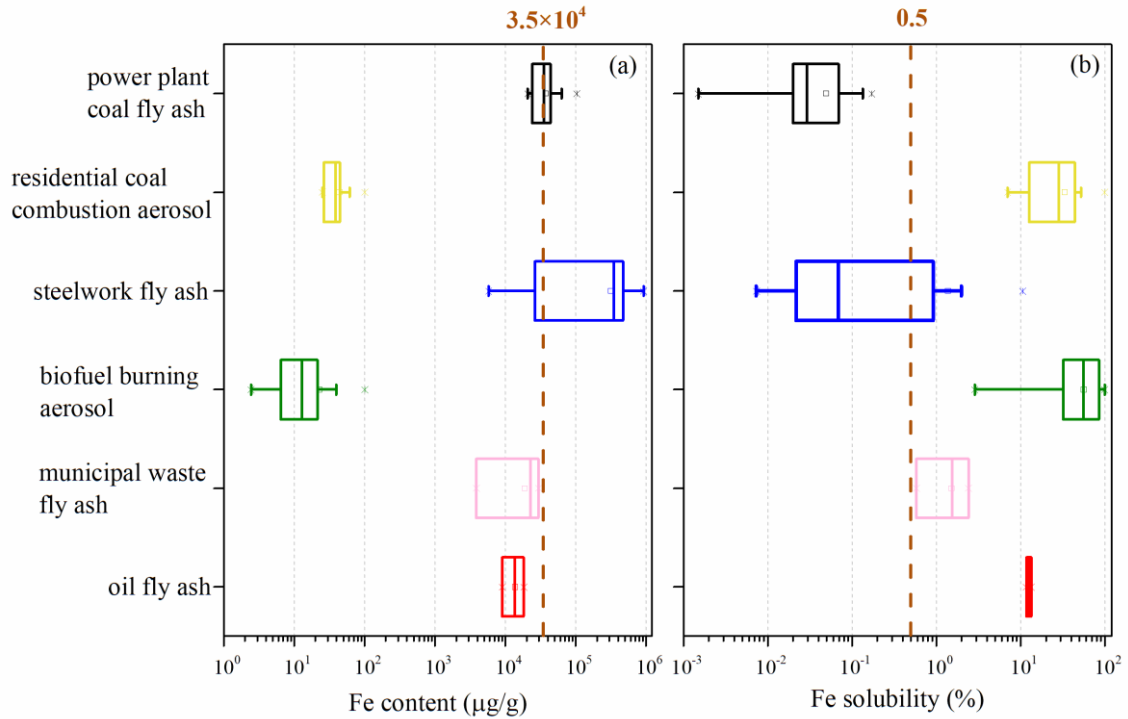
509 Our work considered biofuel burning aerosols for nine types of biofuels, including four
510 types of crop straw, one type of wild grass, and four types of wood. Fe content in biofuel
511 burning aerosols ranged from 0.002 to 0.101 mg/g (Table 2), with average and median values
512 being 0.023 ± 0.026 and 0.013 mg/g, respectively. As shown in Table S6, the average Fe
513 content was determined to be 0.024 ± 0.017 mg/g (n = 3) for PM_{2.5} (Patil et al., 2013), very
514 close to our result; in another study (Hildemann et al., 1991), it was determined to be 0.090
515 mg/g for PM₂ (n = 2), higher than our result. In some other studies, average Fe content were

516 reported to be in the range of 0.162-0.440 mg/g for PM_{2.5} (Alves et al., 2011; Hedberg et al.,
517 2002; Watson et al., 2001; Zhang et al., 2012) and 0.723 ± 0.661 mg/g for PM₁₀ (Schmidl et
518 al., 2008), much higher than our results.

519 Fe content in biofuel burning aerosols showed large variability in different studies, likely
520 due to variations in combustion conditions and biofuel types. For example, metal content in
521 biofuel burning aerosols depended greatly on biofuel types and regions where biofuel was
522 collected (Goncalves et al., 2010), and aerosol particles emitted by wild grass combustion
523 contained larger amounts of metal than wood combustion (Jahn et al., 2021). Modeling studies
524 have used a similar distribution of Fe content between 0.2 and 0.580 mg/g for biofuel burning
525 aerosols (Luo et al., 2008; Rathod et al., 2020).

526 **3.1.5 Fe contents: comparison of anthropogenic and dust Fe**

527 Figure 3a displays Fe content for anthropogenic particles examined in our current study,
528 and the brown dashed line represents the average Fe content of desert dust (35 mg/g) (Taylor
529 and McLennan, 1995). Steelwork fly ash has very high Fe content (median: 346.5 mg/g), about
530 one order of magnitude higher than desert dust. Power plant coal fly ash (median: 35.0 mg/g)
531 has similar Fe content to desert dust. Compared to desert dust, Fe content were around three
532 orders of magnitude lower for residential coal and biofuel burning aerosol (median: 0.038 and
533 0.013 mg/g, respectively). The Fe content was much lower for residential coal and biofuel
534 burning aerosol, likely due to lower combustion temperatures. When combustion occurs at
535 lower temperature, the carbon content of emitted particles is higher; in addition, lower
536 combustion temperature is not sufficient to enable Fe in the fuel to enter emitted particles via
537 volatilization-condensation processes.



538

539 **Figure 3.** Fe content (a) and solubility (b) measured in our work for power plant coal fly ash,
 540 residential coal combustion aerosol, steelwork fly ash, biofuel burning aerosol, municipal
 541 waste fly ash and oil fly ash. The two brown dash lines represent (a) the Fe content (3.5×10^4
 542 $\mu\text{g/g}$) and (b) Fe solubility ($\sim 0.5\%$) for desert dust, respectively.

543 **3.2 Fe solubility by fuel type**

544 **3.2.1 Power plant coal fly ash**

545 Fe solubility in acetate buffer (pH: 4.3) ranged from 0.002% to 0.17% for power plant
 546 coal fly ash (Table 2), with the average and median values being $0.05 \pm 0.05\%$ and 0.03%,
 547 respectively. A few previous studies measured Fe solubility of power plant coal fly ash in
 548 weakly acidic or circumneutral solutions (Table S7). For example, Fe solubility was measured
 549 to be 0.06% in deionized water (Oakes et al., 2012), similar to our result; it was measured to
 550 be 0.2% in dilute sulfuric acid solution (pH: 4.7) (Desboeufs et al., 2005), slightly higher than
 551 our result; the median Fe solubility was determined to be 0.13% in acetate buffer (pH: 4.3) and

552 0.06% in deionized water (Li et al., 2022b), both higher than the median value we obtained.

553 Overall, our work and previous studies suggest that Fe solubility is low in weakly acidic and
554 circumneutral solutions for power plant coal fly ash, with mean or median values around 0.1%.

555 Some studies also measured Fe solubility of power plant coal fly ash in highly acidic
556 solutions and found them to be much higher than those in weakly acidic and circumneutral
557 solutions. For example, Fe solubilities were found to be in the range of 20-25% at pH of 1-2
558 (Chen et al., 2012), 4.2-8.3% at pH of 2 (Fu et al., 2012), and >40% at pH of 2.1 (Baldo et al.,
559 2022). Although Fe solubility measured in strongly acidic solutions may not reflect initial Fe
560 solubility, these studies suggested that acid processing in the emission plume or wider
561 atmosphere could greatly increase Fe solubility for power plant coal fly ash.

562 **3.2.2 Residential coal combustion aerosol**

563 Fe solubility in acetate buffer (pH: 4.3) was determined to range from 7.03% to 100% for
564 residential coal combustion aerosol (Table 2), with the average and median values being 33.30
565 \pm 27.71% and 28.45%, respectively. To our knowledge, no previous study has measured Fe
566 solubility for residential coal combustion aerosol. Compared to power plant coal fly ash, Fe
567 solubility was much higher for residential coal combustion aerosol, and such difference can be
568 attributed to much higher temperature in power plant coal combustion than residential coal
569 combustion. Pyrite (FeS_2) is the major Fe-containing mineral in coal (Deng et al., 2015;
570 Oliveira et al., 2016; Rathod et al., 2020). In low-temperature combustion, pyrite is mainly
571 transformed to Fe sulfate (Bhargava et al., 2009) which has very high Fe solubility; as the
572 temperature increases to >1000 K, Fe sulfate is further transformed to hematite and magnetite
573 which exhibit very low solubility (Hu et al., 2006; Ram et al., 1995; Rathod et al., 2020). A

574 previous study (Rathod et al., 2020) used the relationship between combustion temperature and
575 Fe mineralogy in emitted particles to estimate Fe solubility for different combustion aerosols,
576 and Fe solubility was estimated to be as high as ~32.5% for residential coal combustion
577 aerosols, in good agreement with our experimental results.

578 **3.2.3 Steelwork fly ash**

579 Fe solubility in acetate buffer (pH: 4.3) was determined to range from 0.01% to 10.64%
580 for steelwork fly ash (Table 2), and the average and median values were $1.37 \pm 2.77\%$ and
581 0.07%, respectively. We note that Fe solubility was significantly higher (0.92%-8.59%) for 8
582 samples and very low (<0.5%) for the other 21 samples (Table S3), most of which showed Fe
583 solubility below 0.1%. No previous work has measured Fe solubility for steelwork fly ash. Our
584 experimental results were supported by a modeling study (Rathod et al., 2020) which suggested
585 that the major Fe-containing species in steelwork fly ash were Fe oxides with very low Fe
586 solubilities.

587 **3.2.4 Biofuel burning aerosol**

588 For biofuel burning aerosol, Fe solubility in acetate buffer (pH: 4.3) ranges from 2.86%
589 to 100% with average and median values of $56.07 \pm 30.95\%$ and 55.87%, respectively (Table
590 2). Based on the relationship between combustion temperature and Fe-containing species in
591 emitted aerosols, Fe solubility was previously estimated at 35% for wood burning (i.e., biofuel)
592 aerosol (Rathod et al., 2020), in good agreement with our experimental results.

593 The biofuel examined in our experiment was burnt in a sealed stove and contained no
594 apparent local soil contamination. As such, these results are most representative of domestic
595 biofuel combustion for which the influence of soil-derived Fe can be expected to be negligible.

596 In contrast, wildfires represent dynamic open fire systems that emit aerosol Fe in both fine and
597 coarse fractions (Hamilton et al., 2019). During wildfire combustion, not only is the biofuel
598 (biomass) consumed, but local soils are also entrained into the smoke plumes (Hamilton et al.,
599 2022; Tegler et al., 2023). These soil-derived particles are typically larger (in particle size) and
600 less soluble than their biofuel-derived counterparts (Hamilton et al., 2022), resulting in a larger
601 mass of emitted Fe, albeit with a lower overall Fe solubility. Future studies would benefit from
602 capturing emissions from open burning scenarios to better characterize the properties of
603 wildfire-emitted Fe.

604 **3.2.5 Fe solubilities: comparison of anthropogenic and dust Fe**

605 Figure 3b compares our measured Fe solubility for six types of combustion and
606 anthropogenic particles with that for desert dust. Biofuel burning aerosols (median: 55.87%)
607 and residential coal combustion aerosols (median: 28.45%) exhibited very high Fe solubility.
608 Compared to desert dust, for which Fe solubility is around 0.5% (Chuang et al., 2005; Li et al.,
609 2022b; Ooki et al., 2009; Schroth et al., 2009; Shi et al., 2011), Fe solubility was lower for
610 steelwork fly ash (median: 0.07%) and power plant coal fly ash (median: 0.03%).

611 Overall, Fe solubility in emitted particles was significantly higher for low-temperature
612 combustion (residential and biofuel burning aerosols) than high-temperature combustion
613 (steelwork fly ash and power plant coal fly ash). This is because Fe in emitted particles is
614 mainly highly soluble Fe sulfates for low temperature combustion (Bhargava et al., 2009;
615 Rathod et al., 2020) but Fe oxides with very low solubility for high temperature combustion
616 (Hu et al., 2006; Ram et al., 1995; Rathod et al., 2020).

617 **3.3 Modeling Results**

618 Leveraging new measurements of combustion Fe solubility in residential fuel sources as
619 reported in Sections 3.1 and 3.2, we performed a series of Earth System Model (MIMI)
620 simulations that examined global Fe fluxes in response to modifying anthropogenic Fe
621 solubility parameters at their point of emission. To pair observed solubilities (Table 2) with
622 fuel-types represented in the model, we updated Fe solubility in residential coal burning aerosol
623 from 0.2 to 33% and in biofuel burning aerosol from 10 to 56% at the point of emission (Table
624 3) in both the fine and coarse modes. Biofuel Fe emissions are not limited to residential sources
625 in our Fe inventory. However, lacking updated measurements for industrial biofuel combustion
626 sources, the newly derived solubility parameter (56%) was applied uniformly within the “BIOF”
627 grouping. Waste Fe solubility was added for the central- simulations and was set to 1.5%,
628 following new experimental results provided in Table S7. Smelting Fe solubility was kept at
629 0.03% for all simulations, since new data did not suggest an alternative solubility from what is
630 currently used (Table 3). We ran one additional simulation with updates to industrial sources
631 (PD-IND: industrial coal Fe solubility from 0.2 to 0.05% and oil from 38 to 25%), but observed
632 impacts to global soluble Fe fluxes following these changes were minimal and are accordingly
633 deferred to the Supplement. A description of the fractional solubilities applied to each
634 anthropogenic fuel type within each model simulation is provided in Table 3.

635

636 **Table 3.** Fractional Fe solubilities applied in each model simulation to reflect experimental
637 findings. Bolded rows indicate baseline simulations with no changes made to Fe solubility from
638 previous work using MIMI. To underscore modifications between simulations, a dash (–) is

639 provided where assigned solubility did not differ from the PD-BASE simulation. *Only applied
 640 in the central-residential simulations.

Simulation	Fe solubility modifications by fuel-type (%)					
	Industrial Coal	Residential Coal	Oil	Biofuel	Smelting	*Waste
PD-BASE	0.2	0.2	38	10	0.003	1.5
PD-RESI	-	33	-	-	-	-
PD-BIOF	-	33	-	56	-	-
PD-IND	0.05	33	25	56	-	-
PI-BASE	NA	NA	NA	10	NA	NA
PI-BIOF	NA	NA	NA	56	NA	NA
MID-SSP370-BASE	0.2	0.2	38	10	0.003	NA
MID-SSP370-BIOF	-	33	-	56	0.003	NA
END-SSP370-BASE	0.2	0.2	38	10	0.003	NA
END-SSP370-BIOF	-	33	-	56	0.003	NA

641

642 3.3.1 Impacts on global soluble Fe distribution

643 By applying two new emissions inventories and new solubility parameters for residential
 644 Fe emissions within MIMI, we report a new range of soluble Fe fluxes to the ocean with
 645 regional variations. To isolate the impacts of modifications to each fuel-source, we compared
 646 each model case. When comparing PD-BASE to PD-RESI, the emissions inventory had a
 647 greater influence on the result rather than solubility parameters, per sizeable differences in
 648 residential coal emissions by mass. In the high-residential inventory, global Fe emissions from
 649 residential coal were 460 Gg a⁻¹, exceeding individual emissions from all other fuel types
 650 (industrial coal: 310 Gg a⁻¹; oil: 34 Gg a⁻¹; biofuel: 72 Gg a⁻¹) except for smelting (1350 Gg a⁻¹;
 651 ¹; Table S8). Whereas, in the central- residential inventory, emissions were two orders of
 652 magnitude less at 4.9 Gg a⁻¹ (Table S9), and residential coal sources were the second lowest
 653 contributor to total anthropogenic Fe emissions, next to waste (0.9 Gg a⁻¹). Constraining the Fe

654 content in residential emissions is therefore a more critical knowledge gap to be addressed than
655 constraining the fractional solubility of this source.

656 Accordingly, when applying the central- residential emissions inventory and using the
657 PD-RESI solubility parameters, soluble Fe fluxes to the global ocean only increased by $< 1 \text{ Gg}$
658 a^{-1} (Figure 4; Table 4). Conversely, when applying the high-residential inventory, we saw a 92%
659 increase in anthropogenic soluble Fe fluxes ($+ 33 \text{ Gg a}^{-1}$) to the ocean, translating to a 7%
660 increase in overall soluble Fe fluxes including those from dust and wildfire (Figure 4; Table 4).
661 While the high-residential emissions inventory likely overstates residential coal burning
662 emissions by Fe mass content, the ocean regions most influenced by residential coal burning
663 become apparent, facilitating future research locations of highest interest. Anthropogenic
664 soluble Fe delivery to the Indian Ocean, eastern North Pacific, and parts of the Southern Ocean
665 increased by 300% to 400%, corresponding with a large increases in emissions from China,
666 India, Australia, and South Africa (Figure 4). This follows previous reports of relatively large
667 anthropogenic signals from these regions when compared to global averages (Rathod et al.,
668 2024; Wang et al., 2015).

669

670 **Table 4.** Global present day soluble Fe deposition fluxes in Gg a^{-1} (relative contribution in %)
671 to the ocean for dust, wildfire, and anthropogenic combustion sources. Ranges reflect
672 deposition fluxes between the central- and high-residential emissions inventories (only when
673 variable between cases). Regional fluxes are reported in the Supplement (Table S10).

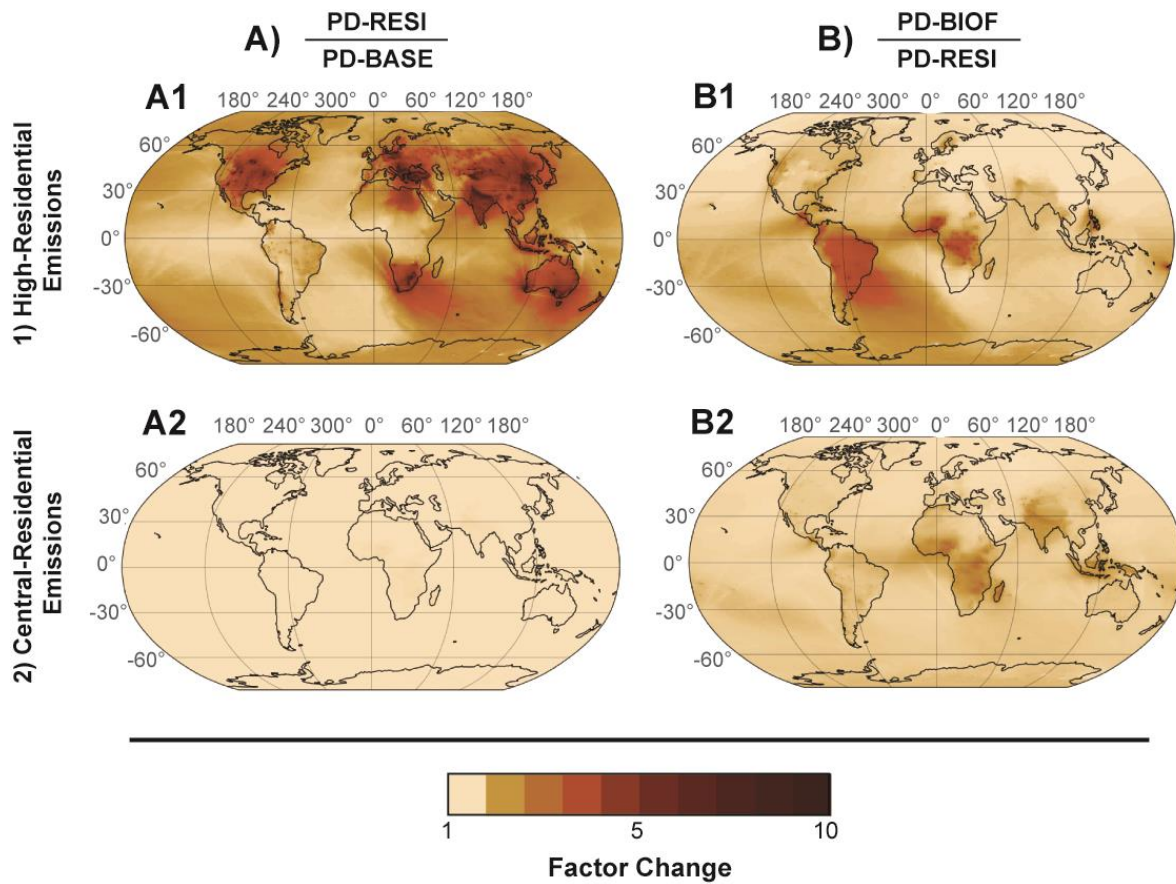
Model Simulation (case)	Dust	Wildfire	Anthropogenic Combustion	All Sources
PD-BASE	305 (83 %)	26 (7 %)	36 (10 %)	367 (100%)

PD-RESI	305 (76-83 %)	26 (7 %)	37-70 (10-17 %)	368-401 (100%)
PD-BIOF	305 (74-81 %)	26 (6-7 %)	49-81 (12-20 %)	375-412 (100%)

674

675 Despite large differences in residential coal emissions between our high- and central-
676 residential inventories, biofuel emissions by mass were nearly identical between inventories
677 (72 and 71 Gg a⁻¹, respectively) because wood burning was already an isolated fuel source in
678 our inventory (Rathod et al., 2020). The small difference is introduced by the separation of
679 waste as a distinct source in the central-residential inventory. When using the high-residential
680 inventory, impacts to soluble Fe fluxes by biofuel were largely overshadowed by residential
681 coal, but using the central- inventory, changes to biofuel parameters in the model controlled
682 increases in soluble Fe fluxes. When we increased biofuel Fe solubility (PD-BIOF), soluble Fe
683 fluxes to the ocean increased by an additional 11 Gg a⁻¹, for a total increase of <1-44 Gg a⁻¹
684 from residential sources (residential coal + biofuel) when compared to PD-BASE (Table 5).
685 Changes to soluble Fe fluxes from biofuel burning were most concentrated across the South
686 Atlantic (Figure 4), likely due to the long range transport of emissions from the Amazon
687 rainforest and across the Congo River basin where biofuel-burning in cook stoves is a common
688 residential practice (García-López et al., 2025; Stoner et al., 2021). It is notable that the signal
689 from central and sub-Saharan Africa impacting the equatorial Atlantic is pronounced using
690 both inventories, but the signal from South America impacting the South Atlantic was much
691 weaker for the central inventory. We report a maximum change in soluble Fe fluxes using the
692 high-residential emissions inventory and PD-BIOF solubility parameters, with deposition to

693 the ocean doubling from 36 to between 49-81 Gg a⁻¹ at the global scale (relative to PD-BASE;
 694 Table 5).



695
 696 **Figure 4.** Relative changes to soluble Fe deposition fluxes using different 1-2) emission
 697 inventories and following modifications to A-B) Fe solubility parameters.

698
 699 With soluble dust-Fe (310 Gg a⁻¹) and wildfire-Fe (30 Gg a⁻¹) fluxes, our different
 700 anthropogenic cases (an additional 13-45 Gg a⁻¹; Table 5) suggest that total soluble Fe fluxes
 701 to the ocean fall between 370 Gg a⁻¹ (PD-BASE) and 380-410 Gg a⁻¹ (PD-BIOF) at the global
 702 scale. These values fall within previous ranges of uncertainty as reported for Fe deposition
 703 fluxes to the ocean (Hamilton et al., 2023; Ito and Miyakawa, 2023), suggesting that solubility
 704 modifications tested herein align with previous Fe constraints within Earth System Models.

705

706 **Table 5.** Global anthropogenic combustion Fe emission and deposition fluxes (Gg a^{-1}) in the
 707 preindustrial (PI), present day (PD), and Future (FU), as simulated by MIMI to two significant
 708 figures. Where ranges are reported reflect use of the central-/high-residential emissions
 709 inventories.

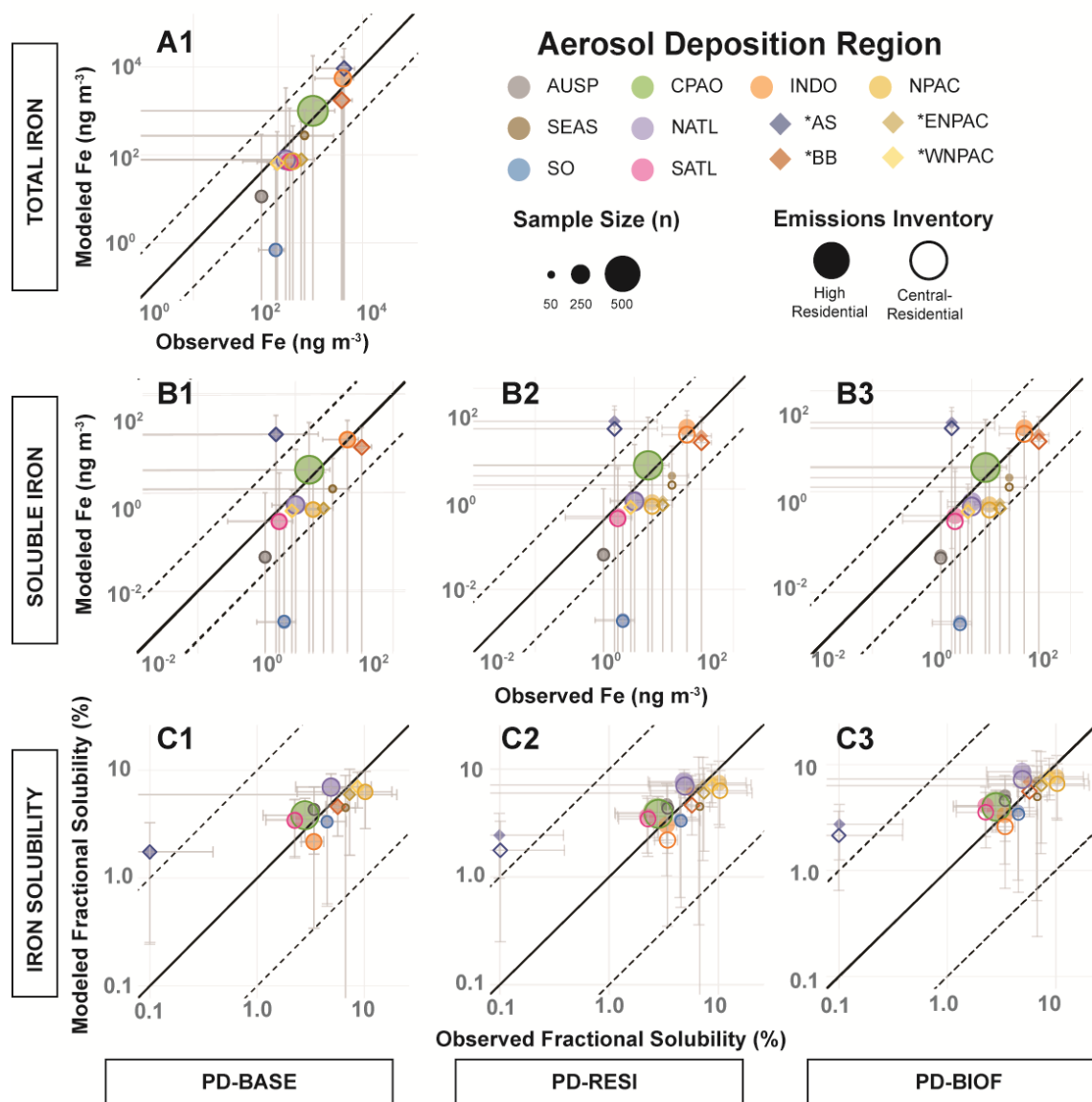
Emission Scenario	Model Simulation (case)	Fe Content	Global Emission	Global Deposition	Deposition to Ocean
PI (1750 CE)	PI-BASE	Total (Soluble)	0.80 (0.08)	0.80 (0.10)	0.30 (0.04)
	PI-BIOF	Total (Soluble)	0.80 (0.44)	0.80 (0.50)	0.30 (0.20)
PD (2010 CE)	PD-BASE	Total (Soluble)	2220 (20-21)	2220 (89-90)	590 (36-40)
	PD-BIOF	Total (Soluble)	2220 (56-170)	2220 (121-270)	590 (46-81)
FU (2050 CE)	MID-SSP370-BASE	Total (Soluble)	2400 (20)	2400 (90)	620 (40)
	MID-SSP370-BIOF	Total (Soluble)	2400 (180)	2400 (250)	620 (80)
FU (2100 CE)	END-SSP370-BASE	Total (Soluble)	1970 (20)	1970 (80)	510 (30)
	END-SSP370-BIOF	Total (Soluble)	1970 (90)	1970 (150)	510 (50)

710

711 3.3.2 Model-observation comparisons of total and soluble Fe concentrations

712 Comparison of modeled surface concentrations with regionally grouped, ship-based
 713 observations revealed generally good agreement between modeled and observed total and
 714 soluble aerosol Fe concentrations for all PD simulations (Figure 5). Modeled total Fe
 715 concentrations were slightly underpredicted when compared to observed values but remained
 716 well within one order of magnitude for each ocean region, with the exception of the Southern
 717 Ocean where total Fe was underestimated by several orders of magnitude. This aligns with

718 previous efforts to model global fluxes of total and soluble aerosol Fe using MIMI v1.0 and
 719 other Earth System Models (Ito and Miyakawa, 2023; Ito et al., 2019; Liu et al., 2024). Current
 720 hypotheses suggest that an Fe source, such as volcanism or mining, is not currently well
 721 represented in models, or alternatively, that limited observations are not representative of Fe
 722 conditions in the airshed of the Southern Ocean (Ito and Miyakawa, 2023; Liu et al., 2024).
 723



724
 725 **Figure 5.** Comparison of modelling and observational data: A) total Fe, B) soluble Fe, and C)
 726 Fe solubility; 1) PD-BASE, 2) PD-RESI, and 3) PD-BIOF. Data were aggregated over time

727 and space as medians. Solid/hollow markers indicate use of the high/central- residential
728 emissions inventories, respectively. Error bars represent spatiotemporal variance within each
729 region. The solid black line indicates a 1-to-1 relationship and the dashed lines represent
730 deviation by ± 1 order of magnitude. Only PD-BASE is shown for total Fe per consistency
731 between cases.

732

733 In each case, soluble Fe regression analyses followed a similar trend to total Fe, wherein
734 modeled averages were slightly lower than observed values and fell within one order of
735 magnitude, apart from the Southern Ocean and Arabian Sea (Figure 5, Table S12). In the PD-
736 RESI and PD-BIOF cases, soluble Fe concentrations increased at the global scale, but the
737 degree varied by region. At the global scale, enhancing residential Fe emissions generally
738 improved model skill for soluble Fe concentrations (Figure 5), resulting in an average increase
739 to modeled soluble Fe concentrations by $0.5 \pm 0.7 \text{ ng m}^{-3}$ within each ocean region (Table S10).
740 For regions most influenced by residential coal burning, the improvement in model skill was
741 slightly higher using the high-emissions inventory, especially for Southeastern Asia (ΔRMSE
742 -0.4), the Bay of Bengal ($\Delta\text{RMSE} = -4.7$), and the eastern North Pacific (Figure 5; Table S12;
743 $\Delta\text{RMSE} = -0.1$). For biofuel burning Fe, the emissions inventory had less of an impact, but
744 enhancing Fe solubility most improved model skill within the South Atlantic Ocean (Figure 5,
745 Table S12 $\Delta\text{RMSE} = -0.05$). Complete summary statistics conveying impacts to model skill for
746 soluble Fe concentrations simulated in each run are provided in the Supplement (Table S12).

747 While the high-residential emissions inventory slightly improved estimates for soluble Fe
748 concentrations, the central-residential emissions inventory performed better when capturing

749 fractional solubility for regions influenced by residential coal (Figure 5). Despite these
750 noticeable variations between cases, ultimately, Fe solubilities calculated by the model were
751 aligned with observations within ± 1 order of magnitude for every region except for the Arabian
752 Sea wherein solubility was overestimated by 1-2 orders of magnitude (Figure 5). In previous
753 MIMI-validation efforts (Hamilton et al., 2019), observational data from the Arabian Sea and
754 Bay of Bengal were aggregated as the Indian Ocean and this result was not flagged. While both
755 basins receive substantial anthropogenic aerosol from India, dust from the Middle East more
756 strongly influences the Arabian Sea, and the Bay of Bengal is more strongly affected by
757 anthropogenic emissions across Southeastern Asia (Bali et al., 2019; Guieu et al., 2019).

758 In general, we found that regions with higher dust inputs more often overshoot
759 measurements when compared to regions with less dust deposition and higher relative impact
760 by anthropogenic emissions (Figure 5; Table S12). In each of our modified solubility cases
761 (PD-RESI and PD-BIOF), Fe solubility for southeastern Asia, the Bay of Bengal, and the North
762 Pacific increased, but the regions heavily impacted by dust remained relatively unchanged
763 (Figure 5).

764 Withstanding source-apportioned measurements of residential coal or biofuel aerosol in
765 our observational dataset, we performed additional model-observation comparisons only using
766 measurements collected in ocean regions downwind of strong residential burning influences.
767 These regions were defined as model-resolution grid cells wherein soluble Fe fluxes increased
768 by 100% or more in the PD-BIOF simulations (Figure 4). However, this reduced median-
769 aggregated observational data points from $n=990$ to $n=25$, limiting statistical capacity to
770 constrain model fluxes. When using the smaller observational dataset, model-observational

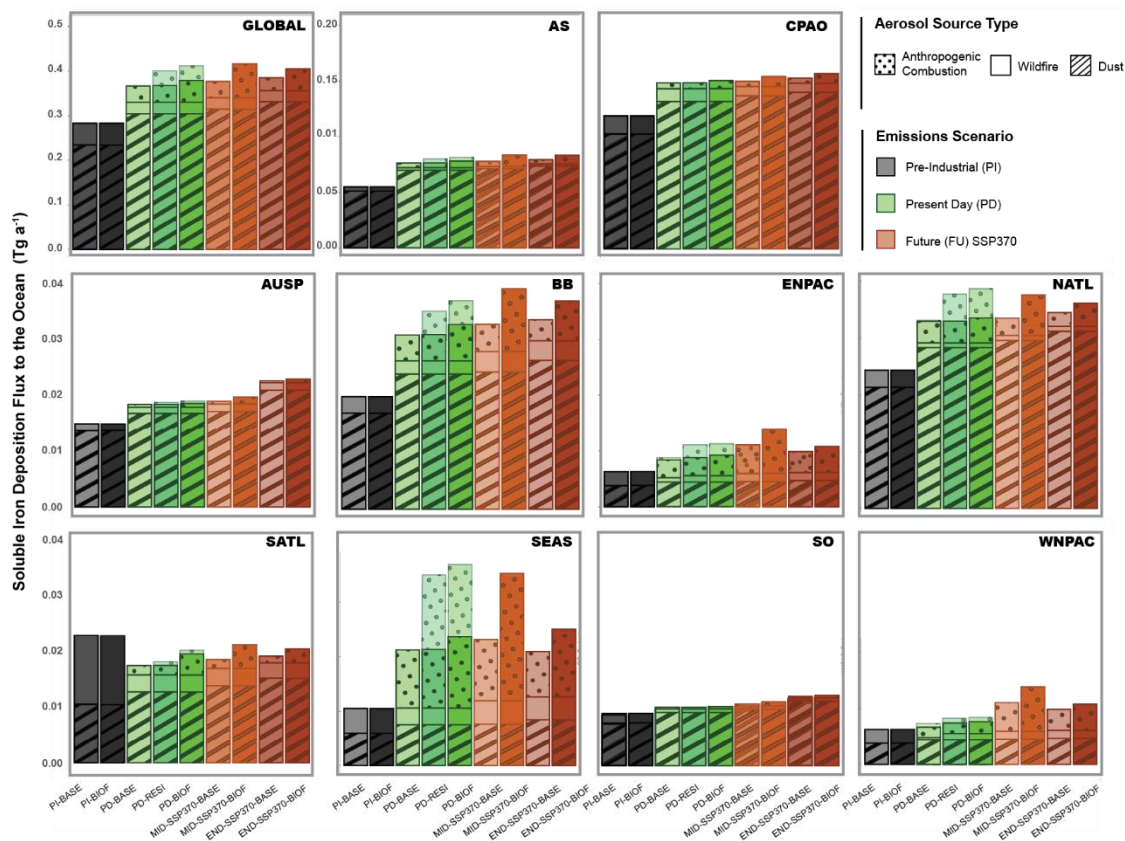
771 comparisons for Total Fe, soluble Fe, and solubility mirrored agreement trends using the larger
772 dataset (Figure 5); those regression analyses are provided in the Supplement (Figures S4-S5).
773 with results from the PD-IND simulation.

774 Anthropogenic soluble Fe fluxes to marine ecosystems increased by 28% to 123% (2% to
775 10% total Fe) when considering all sources of Fe, including dust and biomass burning using
776 the PD-BIOF solubility parameters. For some ocean regions, the increase in soluble Fe fluxes
777 exceeded 150% when using the high-emissions inventory. Despite this, changes to model skill
778 were ultimately modest (Figure 5). The sizeable changes in fluxes with minimal changes to
779 model skill further reinforce the current limitations of ship-based observations in capturing
780 representative soluble Fe fluxes, particularly in the under-sampled Southern Hemisphere and
781 in regions influenced by residential coal and biofuel burning. Future efforts should prioritize
782 expanding the spatial coverage of measurements in these regions to improve model accuracy
783 and understanding of possible anthropogenic influence on remote marine biogeochemistry.

784 **3.3.3 Soluble Fe under PI and FU emission scenarios**

785 PI model simulations serve as a valuable reference point in understanding the specific
786 implications of anthropogenic perturbation to the Earth system. For most regions, soluble Fe
787 fluxes increased between the PI and PD eras (Figure 6), largely attributed to steadily growing
788 anthropogenic combustion emissions and industrial activities over time. Dust and wildfire Fe
789 emissions were also distinct between the PI and PD, due to climatic and land-use change
790 induced feedbacks that have altered global precipitation patterns and dust suspension
791 (Hamilton et al., 2018; Kok et al., 2023; Li et al., 2019; Mahowald et al., 2010). At the global
792 scale, we estimated that current soluble Fe fluxes to marine ecosystems exceed PI fluxes by

793 36-70 Gg a⁻¹, apart from the South Atlantic, wherein soluble Fe fluxes have decreased since
 794 the PI era by 2.8-5.6 Gg a⁻¹ (Figure 6). This decrease was likely attributed to reduction in
 795 wildfire burned area over past decades, particularly in sub-Saharan Africa (Andela et al., 2017;
 796 Jones et al., 2022). Previous work has suggested that wildfire activity during the PI era
 797 exceeded current wildfire regimes at the global scale (Hamilton et al., 2018).



798
 799 **Figure 6.** Deposition fluxes of soluble aerosol Fe to marine ecosystems at the global and
 800 regional scale. Deposition fluxes are source-apportioned (dust, wildfire burning, and
 801 anthropogenic combustion) and provided for each case with distinct solubility parameters. For
 802 the PD, transparent bars represent the high-residential emissions inventory and opaque
 803 represent the central-. Conversely, for the FU, only the high-residential inventory was applied,
 804 and the bars are opaque.

805

806 Regardless of the inventory applied, we observed the largest increases between PI and PD
807 soluble Fe fluxes in Southeastern Asia, the Bay of Bengal, and the North Atlantic. In
808 Southeastern Asia anthropogenic activity has specifically driven, and is projected to drive,
809 changes to future soluble Fe fluxes (Figure 6). This was the only ocean region wherein
810 anthropogenic sources were comparable (central--residential inventory) or exceeded (high-
811 residential inventory) dust sources of soluble Fe (Figure 6). When using the high-residential
812 inventory and the -BIOF solubility parameters, anthropogenic combustion aerosol constituted
813 between 55-72% of soluble Fe fluxes to marine ecosystems in Southeastern Asia, up to a 21%
814 increase from PD-BASE, i.e., based on model solubility parameterizations. Although dust was
815 still the largest source of Fe to the Bay of Bengal and North Atlantic, anthropogenic combustion
816 also strongly influenced soluble Fe delivery in these regions (Figure 6). When using the high-
817 residential emissions inventory with the -BIOF solubility parameters, the relative contribution
818 by anthropogenic emissions doubled in these regions (Table S10).

819 The inventory and solubility parameters used in each case revealed important implications
820 for projected trends of soluble Fe fluxes moving into the second half of the 21st century. Under
821 the SSP370 FU emissions scenario, anthropogenic Fe fluxes were projected to reach their
822 maxima by 2050 for most deposition regions and then decrease to values at or below current
823 PD conditions by 2100 (Figure 6). However, this trend did not hold for regions within the
824 Southern Hemisphere (Australia/South Pacific, Central Pacific/Atlantic, and Southern Ocean)
825 where soluble aerosol Fe fluxes were projected to continually increase through the end of the
826 century (Figure 6). This was not due to direct changes in anthropogenic combustion emissions,

827 but rather due to changes in dust emissions as the primary source of soluble Fe to the Southern
828 Hemisphere (Figure 6).

829 By the end of the century under SSP370, PD-BASE simulations suggest that soluble Fe
830 deposition to global marine systems will increase slightly from 367 to 369 Gg a⁻¹ (~1% increase)
831 by 2100. However, when considering higher residential Fe emissions with higher solubilities
832 representing new upper bound, projected changes to soluble Fe fluxes between PD/MID-
833 SSP370 and MID-SSP370/END-SSP370 was reversed at the global scale and for many regions,
834 with projected decreases by the end of the century (~2% decrease; Figure 6). The projected
835 losses in soluble Fe between the middle and end of the 21st century was especially apparent for
836 the Bay of Bengal (2-6%), the eastern and western North Pacific (12-29%), and across
837 Southeastern Asia (11-41%) (Figure 6), suggesting that various marine ecosystems could face
838 a more significant deviation from current soluble Fe supply than has previously been
839 represented in Earth System Models.

840 Projected changes in soluble Fe fluxes by 2100 under FU emission scenarios, including
841 SSP370, have strong implications for the spatiotemporal distribution of net marine primary
842 productivity, mostly in Fe limited regions. For example, although we demonstrated that
843 anthropogenic emissions most greatly influenced Fe dynamics across Southeastern Asia, it is
844 important to note that primary production in this ocean region is not typically limited by Fe
845 (Bazzani et al., 2023), so ecosystem-level effects by atmospheric Fe are less likely to be
846 observed therein. However, similar to Southeastern Asia, we found that changes to
847 anthropogenic emissions more strongly impacted soluble Fe fluxes in the eastern North Pacific
848 when compared to dust and wildfire sources. Recent work suggests that the atmospheric supply

849 of anthropogenic Fe has already shifted phytoplankton bloom dynamics in the open ocean by
850 accelerating the seasonal uptake of upwelled nitrogen in HNLC regions, including North
851 Pacific (Hawco et al., 2025). Such regions are anticipated to be especially sensitive to changes
852 in anthropogenic Fe given that they are historically limited by trace metals including Fe
853 (Bazzani et al., 2023; Moore et al., 2013; Nishioka and Obata, 2017).

854 In addition to our findings, diverse lines of evidence suggest that half of the soluble Fe
855 flux to the North Pacific comes from Asian anthropogenic sources (Hamilton et al., 2019;
856 Hamilton et al., 2020a; Hawco et al., 2025; Rathod et al., 2020). Li et al. (2024) found that the
857 magnitude of chlorophyll-a response to Fe deposition off the coast of China was lowered by a
858 factor of 4 during COVID-19 in March 2020 when anthropogenic emissions across East Asia
859 were substantially reduced. The authors speculated that a reduction in soluble Fe from
860 anthropogenic activities, either via the primary emission of soluble Fe or via a reduction in Fe
861 solubilization via co-emitted acidic species (e.g., SO_x), resulted in a lessened supply of soluble
862 Fe delivered during the deposition event. Moreover, using Fe isotopes to trace source origins
863 of atmospheric Fe, Hawco et al. (2025) recently showed that the springtime delivery of
864 anthropogenic Fe could be one major factor driving observed seasonal and geographic shifts to
865 the North Pacific transition zone, a highly productive boundary in the North Pacific. Isotopic
866 signatures capable of distinguishing residential coal combustion from other anthropogenic
867 combustion sources have not yet been identified, but our findings suggest that residential coal
868 burning is an especially important source of soluble Fe to the North Pacific and the South China
869 Sea, and across southeastern Asia. Accordingly, we find that projected losses of anthropogenic

870 emissions over the course of this century will most greatly influence nutrient dynamics in these
871 key marine ecosystems.

872 **4 Conclusions**

873 Anthropogenic activity has added a multitude of new aerosol Fe sources to the atmosphere.
874 Understanding how these new sources alter the magnitude and timing of soluble Fe aerosol
875 fluxes to the ocean aids understanding of how human activity is changing marine primary
876 productivity and ocean ecosystem functions within the Anthropocene. However, estimating the
877 contribution of anthropogenic emissions to soluble aerosol Fe fluxes is challenging given the
878 wide variety of sources, each with their own distinct physicochemical profiles. Lack of
879 observational constraints leads to large variation across different modeling studies on the
880 magnitude of the deposition flux from anthropogenic sources. We address some of these
881 uncertainties in this study by measuring the Fe content and solubility of aerosol Fe from several
882 important anthropogenic sources, including a first assessment of the contribution from two
883 major biofuels, namely residential coal and wood. We find that median Fe solubilities vary by
884 greater than three orders of magnitude across fuel types, from 0.03% for power plant coal fly
885 ash to 55.87% for biofuel burning aerosol.

886 To understand the impact of increasing anthropogenic source representations of fractional
887 Fe solubility, we created two new emission inventories that distinguished residential from
888 industrial sources, and further refined Fe solubility parameters for each sources within MIMI,
889 an atmospheric Fe module embedded within the CESM2.

890 At the global scale, we found that current (PD) soluble Fe fluxes to the ocean from
891 anthropogenic sources could exceed current modeled values by somewhere between 28% and

892 123%. This represents a potential increase of over 3 orders of magnitude from the PI when
893 biofuel sources are assumed to be the only source of anthropogenic Fe (Hamilton et al. 2020a).
894 Projected (FU) soluble Fe fluxes from anthropogenic sources remain similar to the PD through
895 to the middle of the century before declining by up to 38% at the end of the century under
896 SSP370.

897 At the regional scale, including residential coal and biofuel burning sources in the model
898 resulted in the most notable impacts for the Bay of Bengal, across Southeast Asia, and
899 throughout the North Pacific and North Atlantic (i.e., regions strongly influenced by nearby
900 continental anthropogenic activity). However, these regions are generally under-sampled in
901 terms of shipborne aerosol Fe observations, and therefore, to reduce the largest source of
902 uncertainty, more measurements are needed in regions downwind of residential Fe sources to
903 better constrain the contribution of human activity on global biogeochemical cycles.

904

905 **Data availability.**

906 Experimental data can be found in the manuscript or the supplement, or are available at
907 <https://doi.org/10.5281/zenodo.17766592> (Li, 2025). Modeling output data, coding scripts are
908 available at <https://github.com/haleyplaas/CombustionFe>; emission inventories are available at
909 <https://doi.org/10.5281/zenodo.17903186>.

910 **Competing interests.**

911 At least one of the (co-)authors is a member of the editorial board of Atmospheric
912 Chemistry and Physics.

913 **Author contribution.**

914 MT initiated this study; MT and DSH designed this study and secured funding resources;
915 RL, YZ, YC and TZ conducted experimental work; HEP, SR and DSH conducted modeling
916 work; YY provided key samples used in this work and contributed to data analysis; RL and
917 HEP analyzed the results; RL, HEP, DSH and MT wrote the manuscript; all the authors
918 reviewed and approved the manuscript.

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923

924

925 **References**

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