

# 1 **Meta-analytical insights into organic matter enrichment in the surface** 2 **microlayer**

3 Amavi N. Silva<sup>1</sup>, Surandokht Nikzad<sup>2,3</sup>, Theresa Barthelmeß<sup>1</sup>, Anja Engel<sup>1</sup>, Hartmut Herrmann<sup>4</sup>, Manuela  
4 van Pinxteren<sup>4</sup>, Kai Wirtz<sup>3,5</sup>, Oliver Wurl<sup>6</sup> and Markus Schartau<sup>1</sup>

5 <sup>1</sup>GEOMAR Helmholtz Centre for Ocean Research Kiel, Kiel, 24143, Germany

6 <sup>2</sup>Trent University, Peterborough, ON K9L 0G2, Canada

7 <sup>3</sup>Institute of Coastal Systems, Helmholtz-Zentrum Hereon, Geesthacht, 21502, Germany

8 <sup>4</sup>Leibniz Institute for Tropospheric Research (TROPOS), Atmospheric Chemistry Department (ACD), Leipzig, 04318,  
9 Germany

10 <sup>5</sup>Christian-Albrechts University of Kiel, Kiel, 24148, Germany

11 <sup>6</sup>Carl von Ossietzky University of Oldenburg, Oldenburg, 26129, Germany

12 *Correspondence to:* Amavi N. Silva ([asilva@geomar.de](mailto:asilva@geomar.de)), Markus Schartau ([mschartau@geomar.de](mailto:mschartau@geomar.de))

13 **Abstract.** The surface microlayer (SML), the uppermost ~1 mm water layer at the air-water interface, plays a critical role in  
14 mediating Earth system processes, yet current knowledge of its composition and organic matter enrichment remains scattered  
15 across disciplines. Here, we present the first known meta-analysis of SML studies that quantitatively assesses the distributional  
16 characteristics of selected organic compounds, including organic carbon and nitrogen, amino acids, fatty acids, transparent  
17 exopolymer particles, carbohydrates, lipids and proteins, through probability density estimates, central tendency metrics and  
18 correlations analyses. Our results confirm a preferential enrichment of nitrogen-enriched, particulate organic matter in the  
19 SML, while also highlighting the significance of surfactant-specific factors that govern selective enrichment in the SML. We  
20 find that enrichment patterns can vary systematically with environmental and methodological conditions, underscoring the  
21 need to account for such influences when interpreting observations and developing SML-based models. We provide the full  
22 range of typical EF values for the studied compounds, offering a clear reference for assessing whether new measurements are  
23 typical or extreme. While delving into the ability of EFs to reflect organic matter partitioning in the SML, we also critically  
24 examine their limitations in capturing trophic variability and suggest that EF-based assessments be complemented with metrics  
25 that remove background variability from underlying water concentrations, enabling more accurate interpretations of true SML  
26 enrichment and informing future modelling efforts. Additionally, our meta-analysis demonstrates that logarithmic data  
27 transformations and robust central tendency estimates outperform traditional linear-scale approaches, providing more accurate  
28 and reliable SML enrichment estimates.

## 29 **1 Introduction**

30 Approximately 70% of the Earth's surface is covered by a hydrated gelatinous 'skin' known as the surface microlayer  
31 (hereafter referred to as 'SML'; note that while this term is commonly used to denote the sea surface microlayer, in this study

32 it refers to the surface microlayer in both marine and freshwater systems), which has an operationally defined thickness  
33 typically ranging from 1 – 1000  $\mu\text{m}$ , depending on the sampling method used (i.e., screen, plate, drum: Astrahan *et al.*, 2016;  
34 Hunter, 1980; Liss and Duce, 1997, 1997; Wurl *et al.*, 2009). Situated between the surface waters of all natural water bodies  
35 and the atmosphere, this uppermost multi-component layer (Astrahan *et al.*, 2016; Carlucci *et al.*, 1985; Cunliffe *et al.*, 2013)  
36 creates a unique microhabitat, mainly consisting of neuston (i.e., living communities in the SML), a relatively enriched  
37 complex of organic compounds and strong physico-chemical gradients (Cunliffe *et al.*, 2013; Dietz *et al.*, 1976; Engel and  
38 Galgani, 2016; Hunter and Liss, 1977). The formation and the composition of the SML are governed by a number of biological,  
39 physical and chemical drivers that interact under varying complex environmental conditions and time scales. As a result, the  
40 SML dynamics play a pivotal role in a range of environmental processes such as air-water gas exchange, heat transfer across  
41 boundary layers, biogeochemical cycling, microbial interactions and distribution of pollutants (e.g., Engel *et al.*, 2017; Frew,  
42 1997; Liss and Duce, 1997; Upstill-Goddard, 2006). Therefore, continued investigation of the compositional heterogeneity of  
43 the SML and of the processes therein is crucial to gain deeper insights into its role in ocean biogeochemistry and its potential  
44 climate interactions.

45 The SML is shaped by physical forces: surface tension provides structural stability at the air-water interface (Liss and Duce,  
46 1997), while diffusive fluxes, bubble scavenging and the upward transport of buoyant particles deliver material from  
47 underlying waters (hereafter referred to as ‘ULW’; Baastrup-Spohr and Staehr, 2009; Chen *et al.*, 2016; Joux *et al.*, 2006;  
48 Obernosterer *et al.*, 2005). In addition, wet and dry atmospheric deposition as well as *in situ* production and degradation also  
49 lead to concentration changes in the SML (Astrahan *et al.*, 2016; Kuznetsova *et al.*, 2004; Milinković *et al.*, 2022). Within the  
50 SML, biological and chemical processes continuously transform compounds between dissolved and particulate forms (Liss  
51 and Duce, 1997), further contributing to its characteristic enrichment relative to the ULW (e.g., Baastrup-Spohr and Staehr,  
52 2009; Gao *et al.*, 2012; Gašparović *et al.*, 2007; Liss and Duce, 1997; Marty and Saliot, 1976; Yang, 1999).

53 Many compounds present in the SML are surface active and are collectively known as ‘surface-active-agents’ or ‘surfactants’  
54 (Maki and Hermansson, 2020; Wurl and Holmes, 2008). Surfactants tend to adsorb at the air-water interface (Wurl *et al.*, 2009)  
55 due to their amphiphilic nature (i.e., presence of both hydrophobic and hydrophilic structural parts; e.g., Marty and Saliot,  
56 1976) and form interfacial films. This leads to modifications of the physico-chemical characteristics of the sea surface, most  
57 notably surface tension, elasticity and viscosity, which alter momentum transfer, micro-scale wave breaking, damping of  
58 capillary waves, ultimately affecting air-sea gas exchange (McKenna and McGillis, 2004; Pereira *et al.*, 2016). Selective  
59 enrichment of surfactants in the SML is strongly influenced by phytoneuston exudation and grazing processes (Kujawinski *et al.*,  
60 2002; Žutić *et al.*, 1981), which release carbohydrates that constitute a major fraction of naturally occurring biosurfactants  
61 (Myklestad, 1995; Penna, 1999). Blooms facilitate the accumulation of large hydrophilic polysaccharides, which can bind to  
62 hydrophobic groups and thereby acquire surfactant properties (Wurl *et al.*, 2011). Surfactant distribution is further shaped by  
63 microbial activity (Hunter and Liss, 1977; Kurata *et al.*, 2016); Baceterioneuston is predominantly lipolytic and proteolytic,

64 breaking down organic matter (OM) into lipids and proteins (polymers of amino acids; Carlucci *et al.*, 1985; Kjelleberg *et al.*,  
65 1976; Sieburth *et al.*, 1976), both of which represent abundant biosurfactant pools in the SML (Brinis *et al.*, 2004; Marty and  
66 Saliot, 1976). However, carbohydrates and polysaccharides also constitute major, rapidly utilized substrates for heterotrophic  
67 bacteria in the SML (Harvey *et al.*, 1995; Penezić *et al.*, 2022).

68 Surfactants have been categorized according to their solubility into dry and wet surfactants, of which the more insoluble  
69 fraction tends to establish as a monolayer at the surface (e.g., phospholipid-like material; Frka *et al.*, 2012), while the adsorption  
70 of the latter fraction (more soluble; e.g., proteins and carbohydrates) is governed by concentration-driven equilibria  
71 (Asmussen-Schäfer *et al.*, 2026; Laß and Friedrichs, 2011). Nonetheless, the natural soluble surfactant pool frequently reaches  
72 a threshold beyond which monolayer-like surfactant coverage of the air–sea interface is observed (Asmussen-Schäfer *et al.*,  
73 2026). In addition to their chemical composition, surfactants also vary in their size: Colloidal and particulate organic matter  
74 accumulated in the SML further provide substrates to bacterioneuston, thereby helping to stabilize the surface films (Sieburth,  
75 1983). The contribution from the particulate pool to the SML’s surface activity is estimated to range from 10% to 55%  
76 (Gašparović and Čosović, 2003). Furthermore, sticky microgels, like transparent exopolymer particles (TEP) that originate  
77 from bacteria and phytoplankton (Alldredge *et al.*, 1993), are also found in the SML. Such gel-like particles can form through  
78 the coagulation of dissolved polysaccharides (Engel *et al.*, 2004; Mari and Burd, 1998; Schartau *et al.*, 2007), and are capable  
79 of incorporating other compounds into a cohesive matrix (Cunliffe *et al.*, 2009; Sieburth, 1983; Wurl and Holmes, 2008),  
80 thereby enhancing the structural integrity of surface films (Cunliffe and Murrell, 2009). When the SML becomes highly  
81 concentrated in surfactants, these films transform into thick surface slicks that are visible to the naked eye (Liss and Duce,  
82 1997). The extent to which OM-driven changes in SML surfactant composition alter air-sea gas exchange remains to be fully  
83 understood (Pogorzelski *et al.*, 2006). In addition, inorganic ions, which do not preferentially adsorb at the air-water interface,  
84 can be also present in the SML due to passive upward transport (Knipping *et al.*, 2000; Petersen *et al.*, 2004).

85 Liss and Duce (1997) and Pereira *et al.*, (2018) argue that the SML can restrict diffusive fluxes across the air-sea interface,  
86 substantially contributing to reduced rates of ocean-atmosphere gas exchange. Surfactants can impact air-sea gas exchange of  
87 greenhouse gases such as carbon dioxide (CO<sub>2</sub>), methane (CH<sub>4</sub>), nitrous oxide (N<sub>2</sub>O) and dimethyl sulfide (DMS) (Frew, 1997;  
88 Upstill-Goddard, 2006). Asher (1997), from laboratory measurements, and Tsai and Liu (2003), from global ocean  
89 observations, estimate a reduction of annual net CO<sub>2</sub> flux by ~20% – 50% due to the presence of the SML, while Wurl *et al.*,  
90 (2016), from *in situ* measurements, propose that this decrease can be ~15%. Barthelmeß *et al.* (2021) observed that, in a newly  
91 upwelled filament off Mauritania, surfactants can suppress CO<sub>2</sub> gas exchange by 12%. Both lab- and field-based experiments  
92 find that natural slicks can reduce air-sea gas exchange by 50 – 60% (Goldman *et al.*, 1988; Salter *et al.*, 2011; Mustaffa *et al.*,  
93 2020), causing the SML to drive an overall reduction of 19% in the CO<sub>2</sub> fluxes, as shown by *in situ* observations (Mustaffa *et al.*  
94 *et al.*, 2020). Supporting earlier findings of Springer and Pigford (1970), McKenna and McGillis (2004) and Sabbaghzadeh *et al.*  
95 (2017), who raised concerns about the impact of the SML’s surfactants on uncertainties in air-sea gas exchange models,

96 Mustaffa *et al.* (2020) further argue that conventional wind-based models miscalculate CO<sub>2</sub> exchange up to 20% in areas with  
97 high surfactant concentrations. Moreover, Kock *et al.* (2012) find that, in the eastern tropical North Atlantic region, offsets  
98 between air-sea and diapycnal N<sub>2</sub>O fluxes could be explained when surfactant effects were introduced to gas exchange models.  
99 Work of Goldman *et al.*, (1988) find that surfactants in the SML can also suppress air-sea gas exchange of oxygen (O<sub>2</sub>).  
100 Disparities in these studies emphasize the significance of accurately assessing the characteristics of the SML and its processes,  
101 as well as integrating this knowledge into climate relevant ocean-atmosphere models (Milinković *et al.*, 2022) in order to  
102 reduce uncertainties in global gas flux estimations, particularly given that SML is seldom included in gas exchange models  
103 (Cen-Lin and Tzung-May, 2013; Engel *et al.*, 2017).

104 Although the composition and the concentration of compounds within the SML are thought to be strongly correlated with those  
105 of the ULW (Basstrup-Spohr and Staehr, 2009; Chen *et al.*, 2016; Joux *et al.*, 2006; Kuznetsova *et al.*, 2004), certain substances  
106 are selectively accumulated at the air-water interface, leading to a pronounced enrichment in the SML. Several studies,  
107 including Carlucci *et al.* (1985), Henrichs and Williams (1985), Kuznetsova and Lee, (2002) and Reinthaler *et al.* (2008), find  
108 stronger enrichment of particulate fractions and nitrogen-based compounds compared to dissolved organic carbon. The  
109 accumulation of these specific compounds in the SML relative to the ULW is often described by the ‘Enrichment Factor’  
110 (hereafter referred to as ‘EF’). The EF of a compound ‘x’ is given by the following concentration ratio:

$$111 \quad EF \text{ of } x = \frac{\text{Concentration of } x \text{ in SML}}{\text{Concentration of } x \text{ in ULW}} \quad (1)$$

112 According to this equation, when the concentration of x is higher in the SML than in the ULW, the EF value rises above 1;  
113 when it is lower, the EF drops below 1, as discussed in Carlson (1983) and Garabetian *et al.* (1993). However, previous studies  
114 report substantial enrichment variability in the SML across environments, compound classes and spatio-temporal scales. For  
115 instance, non-slick areas where microbial degradation processes are dominant can also demonstrate higher EF values,  
116 resembling those found in slick conditions (e.g., Bastrup-Spohr and Staehr, 2009). In contrast, some lakes appear to exhibit  
117 weak SML enrichment even under eutrophic ULW conditions when the waters are concentrated by autochthonous OM (i.e.,  
118 originate within the same ecosystem they are found) that show a lower affinity to the air-water interface (Hillbricht-Ilkowska  
119 and Kostrzewska-Szlakowska, 2004). Freshwater SML tends to be more enriched with organic carbon and nitrogen, total  
120 phosphorous, ammonia and phosphate ions (Knulst *et al.*, 1997; Münster *et al.*, 1998; Södergren, 1987), whereas in marine  
121 environments, carbohydrates, lipids, proteins and amino acids tend to be more enriched (Liss and Duce, 1997). Concentration  
122 variability of the SML can be significantly larger than that of the ULW (Reinthaler *et al.*, 2008) though, in some occasions the  
123 two layers show similar variability (Carlson, 1983). Likewise, the extent to which SML composition mirrors the ULW also  
124 varies, with some studies observing tight coupling (e.g., Chen *et al.*, 2016; Joux *et al.*, 2006) and others reporting marked  
125 decoupling linked to different mineralization rates or adsorption dynamics (e.g., Kuznetsova *et al.*, 2004). Differing surface

126 activities (i.e., tendency of a substance to accumulate at the interface) of organic compounds is considered a major driver of  
127 these transfer dynamics between the SML and the ULW (Engel *et al.*, 2017). Environmental factors further influence SML  
128 enrichment, yet their influence remains inconsistent and unresolved (e.g., Baastrup-Spohr and Staehr, 2009; Carlson, 1983;  
129 Reinthaler *et al.*, 2008; Sabbaghzadeh *et al.*, 2017). Collectively, these heterogeneous findings highlight the complexity of SML  
130 enrichment processes and the need for systematic-cross-study evaluations. These aspects and their implications are further  
131 discussed in the Discussion.

132 Overall, the diversity of reported findings highlights the need for a more holistic view of the applicability of EF as a valid and  
133 meaningful indicator of compounds enriched in the SML. To address this, we adopted a meta-analysis of existing SML-studies,  
134 and conducted a comprehensive analysis to (1) assess OM enrichment in the SML, (2) review current EF estimates and (3)  
135 investigate the relevance of EF values as accurate indicators of OM enrichment. The data collection presented here covers  
136 mass concentrations of OM compounds and does not include measurements of surface activities or effects on the physico-  
137 chemical properties of the uppermost monolayer of the SML. The primary objective is to provide an overview and specific  
138 insights into OM compounds that can accumulate within the SML and potentially be linked to biogeochemical processes  
139 occurring in the ULW. Accordingly, surfactant measurements of surface activities that have been converted into equivalent  
140 surfactant concentrations, such as those expressed as Triton X-100 equivalents, are not considered here. Ultimately, this data  
141 compilation, together with the knowledge derived from its initial meta-analysis, is intended to establish a robust foundation  
142 for subsequent studies that may support future modelling efforts linking biological processes to functions of the SML and their  
143 implications for biogeochemistry and climate.

144

## 145 **2 Methodology**

146 The work presented here synthesizes findings from multiple studies on the SML and employs a quantitative meta-analysis.  
147 Meta-analyses provide an essential means of extracting robust and generalizable conclusions by integrating results from  
148 fragmented bodies of literature. Such systematic reviews can provide a more precise and accurate understanding of overarching  
149 trends, even when individual studies report inconsistent results (Crocetti, 2016). Mengist *et al.* (2020) highlight the importance  
150 of meta-analyses by stating that “*Systematic reviews with meta-analysis represent the gold standard for conducting reliable*  
151 *and transparent reviews of literature.*” In fields such as SML research, where methodological diversity is high and  
152 environmental variability is inherent, meta-analytical approaches are invaluable in identifying coherent trends and key  
153 constraints.

## 154 2.1 Data collection and compilation

155 The primary dataset consists of 2055 data points, extracted from 31 peer-reviewed publications (hereafter referred to as  
156 ‘reference studies’) identified through a comprehensive and systematic literature search of scholarly articles published between  
157 1967 and 2022. These studies were identified through a structured Google Scholar search conducted between February 2025  
158 to April 2025. Search terms included combinations of descriptors related to the surface microlayer (e.g., ‘sea surface  
159 microlayer’, ‘SML’), enrichment terminology (e.g., ‘enrichment’, ‘enrichment factor’), and compound specific keywords (e.g.,  
160 ‘surfactants’, ‘organic carbon’, ‘organic nitrogen’, ‘TEP’, ‘Amino acids’, and other organic matter classes). Additional relevant  
161 publications were identified through reference lists of retrieved papers. Only datasets providing extractable numerical values  
162 were retained.

163 From these studies, directly measured mass concentration data were extracted from simultaneously collected SML and ULW  
164 samples (hereafter referred to as  $[C]_{\text{SML}}$  and  $[C]_{\text{ULW}}$ , respectively) for twelve different observational types of organic  
165 compounds (hereafter known as ‘target compounds’): total organic carbon (TOC expressed in  $\text{mg L}^{-1}$ ), particulate organic  
166 carbon (POC in  $\text{mg L}^{-1}$ ), dissolved organic carbon (DOC in  $\text{mg L}^{-1}$ ), total organic nitrogen (TON in  $\text{mg L}^{-1}$ ), particulate organic  
167 nitrogen (PON in  $\text{mg L}^{-1}$ ), dissolved organic nitrogen (DON in  $\text{mg L}^{-1}$ ), amino acids (AA in  $\mu\text{mol L}^{-1}$ ), fatty acids (FA in  $\mu\text{g}$   
168  $\text{L}^{-1}$ ), transparent exopolymer particles (TEP in  $\mu\text{g Xeq L}^{-1}$ ), carbohydrates (CHO in  $\mu\text{mol L}^{-1}$ ), lipids (in  $\mu\text{mol L}^{-1}$ ) and proteins  
169 (in  $\mu\text{mol L}^{-1}$ ). TOC pool includes all forms of organic carbon, thus comprising both POC and DOC. Similarly, the TON pool  
170 combines both PON and DON. In general, the particulate pool constitutes a minor fraction of the total pool. The major classes  
171 of biopolymers are proteins, CHO and lipids, with AA serving as the monomers of proteins. Depending on the elemental  
172 composition of these biopolymers, they contribute to both, the organic carbon and/or organic nitrogen pool. While the ratio of  
173 these biopolymers is higher in the particulate pool, it usually declines to only a few percent in the dissolved pool. TEP is  
174 composed of polysaccharides (i.e., CHO) with a major fraction contributing to POC, while a minor fraction exists at the  
175 interface between the dissolved and particulate phases (Verdugo *et al.*, 2004). These compounds were selected as they represent  
176 major carbon and nitrogen pools in the SML, are widely reported across marine and freshwater systems, and are sufficiently  
177 represented in the literature to support a robust meta-analytical assessment.

178 The EF values for these target compounds were systematically calculated from corresponding  $[C]_{\text{SML}} - [C]_{\text{ULW}}$  pairs, using Eq.  
179 (1). In this study,  $[C]_{\text{SML}}$ ,  $[C]_{\text{ULW}}$  and EF data are collectively referred to as ‘primary data’. Auxiliary information associated  
180 with the primary data, (i.e., sampling factors and environmental variables) were also extracted when reported and are referred  
181 to as ‘secondary data’. All analyses were performed using the complete primary data set, independent of whether secondary  
182 data were available. The secondary data were summarized only to illustrate existing research gaps in SML studies. All the data  
183 were collected either (1) directly from the source when presented, or else (2) through digitization of graphs and plots using  
184 PlotDigitizer (<https://plotdigitizer.com>) and GraphClick v3.0 (<https://graphclick.en.softonic.com/mac>). To estimate

185 digitization uncertainty, TOC data (40 datapoints) from Baastrup-Spohr and Staehr (2009) were digitized five times (200  
186 values in total). The standard deviation of repeated measurements was calculated for each point and expressed relative to its  
187 average. Across all points, the median relative uncertainty was 0.29%, indicating that digitization introduced minimal error.  
188 The resulting compiled database is herein referred to as ‘*Surface Microlayer Organic Matter Global Data Collection*’ (SML-  
189 OM). Supplementary Table S1 provides an overview of the reference studies on which SML-OM is based.

## 190 **2.2 Statistical analyses**

191 Given that the SML-OM ranges over several orders of magnitudes, when the dataset is handled in linear-space (i.e. in its  
192 original form), higher values dominate and overshadow the features associated with lower values (Feenstra, 2006). These  
193 potential limitations of linear scaling were reduced by transforming our primary data into their logarithmic ( $\log_{10}$ ) counterparts.  
194 Hereafter, the term ‘linear’ refers to the original, untransformed data, while the term ‘log’ stands for their logarithmic  
195 equivalents. The following sections describe the subsequent analyses conducted in our work.

### 196 **2.2.1 Probability distributions**

197 Making inferences based on ratios such as EFs requires careful consideration, as changes in the numerator and the denominator  
198 often affect these ratios asymmetrically (Keene, 1995). In the context of this study, while reductions in  $[C]_{ULW}$  can lead to  
199 unusually high EF values that can approach infinity (i.e., stretched towards higher values), increases in  $[C]_{ULW}$  may produce  
200 EFs decreasing down to 0 (i.e., compressed towards lower values). This results in distributions that significantly deviate from  
201 Gaussian (i.e. normally distributed) shape. Therefore, distributional characteristics of the primary data were examined through  
202 probability distributions.

203 Probability density functions (hereafter referred to as ‘PDF’) of the EF values were examined by applying non-parametric  
204 Kernel Density Estimates (hereafter referred to as ‘KDE’; Parzen, 1962; Silverman, 1986; Wegman, 1972). KDE employs a  
205 normalized weighting function – known as ‘Gaussian kernel’ – which is centered at each datapoint. The sum of these kernels  
206 produces a smooth and continuous PDF that fits the underlying data. Selection of the width of a kernel – known as ‘bandwidth’  
207 – is an integral part of the KDE approach, as bandwidths too small or too large lead to overfitting and underfitting of data,  
208 respectively, failing to capture the true patterns in distributions. Following this, optimal bandwidths for linear KDEs were  
209 computed based on Härdle *et al.* (2004). For log KDEs, a fixed bandwidth was applied. **Log transformations, unlike the linear  
210 scale, produce similar distributions with comparable spreads across variables, allowing a single fixed bandwidth to produce  
211 stable and consistent smoothing for all data.**

212 Robustness of the KDE method decreases at low sample size. Since the SML-OM contains variables with sample sizes as low  
213 as 16 (for proteins), a bootstrap resampling approach was adopted where 67% of the original data (i.e., 2/3 of the sample) were  
214 randomly subsampled. This proportion balances the need for sufficient data to generate stable KDEs while still introducing

215 variability for robustness testing. By allowing consistent treatment across all data types, this approach maintains comparable  
216 KDE bandwidth behavior among subsamples. An individual KDE was generated at each iteration. The process was repeated  
217 1000 times, each time with a different random subsample, generating a set of KDE. These were then averaged to produce an  
218 ensemble mean, from which the final PDFs were derived.

219 Additionally, cumulative distribution functions (hereafter referred to as ‘CDF’) were determined for  $[C]_{\text{SML}}$  and  $[C]_{\text{ULW}}$  from  
220 the ensemble means of the bootstrapped KDEs. Appendix A provides further information on the KDE method.

## 221 2.2.2 Summarization, comparison and correlation estimates of distributions

222 For describing, comparing and relating the resulting PDFs and CDFs, we used standard statistical measures. Their  
223 mathematical expressions are given in Appendix B.

224 (1) To describe the central tendencies, mode (hereafter referred to as ‘ $x_m$ ’), median (hereafter referred to as ‘ $\tilde{x}$ ’),  
225 arithmetic mean (hereafter referred to as ‘ $\bar{x}_a$ ’) and geometric mean (hereafter referred to as ‘ $\bar{x}_g$ ’), were computed.

226 (2) The values at 5<sup>th</sup> and 95<sup>th</sup> percentiles of each distribution (hereafter referred to as ‘upper threshold: UT’ and ‘lower  
227 threshold: LT’, respectively) were also estimated in order to determine their central 90% range (i.e., degree of spread).

228 (3) To numerically compare the  $[C]_{\text{SML}}$  and  $[C]_{\text{ULW}}$ , Integrated Quadratic Distance (Hereafter known as ‘IQD’) values of  
229 their CDFs were approximated based on Eq. (B3), which measure how different the two distributions are with regard  
230 to symmetry and multimodality.

231 (4) To investigate and quantify potential relationships between  $[C]_{\text{SML}}$  and  $[C]_{\text{ULW}}$  of each target compound, their linear  
232 correlation was analysed by employing both parametric Pearson and non-parametric Spearman’s tests (both methods  
233 were applied for cross-validation purposes; agreement between the two correlation coefficient values increases the  
234 confidence in the robustness of the observed relationship).

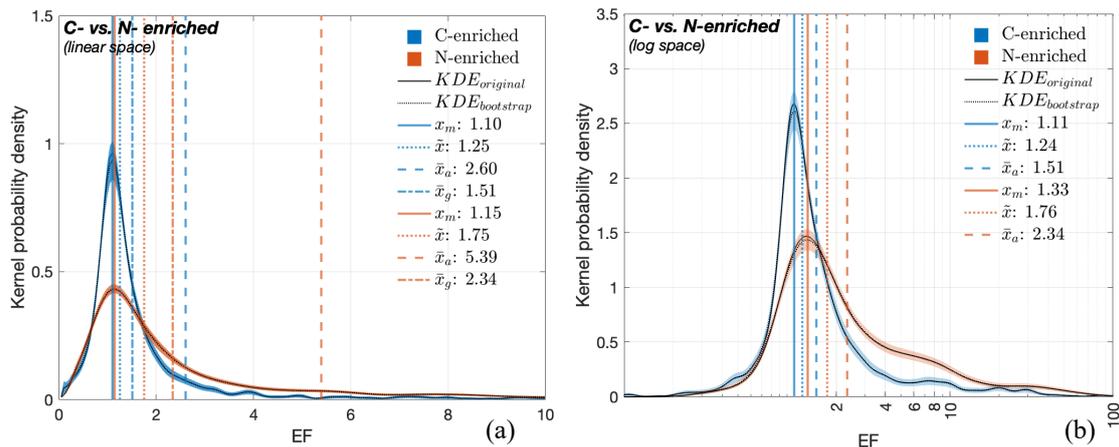
## 235 3 Results

236 Unless otherwise stated, all analyses were performed on log scale. Nevertheless, to avoid potential misinterpretation of log  
237 scales in data presentation, primarily due to their limited readability among non-expert audiences (e.g., Menge *et al.*, 2018),  
238 all results are presented on linear scale.

### 239 3.1 Characterizing EF distributions

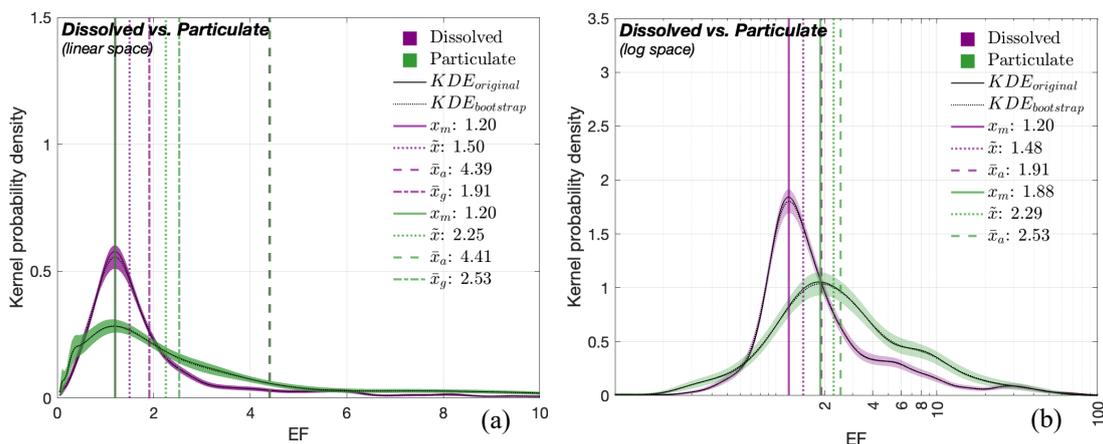
240 Figure 1 compares the KDE-derived PDFs of the EF values for the carbon-enriched (in blue) and nitrogen-enriched (in orange)  
241 organic compounds (Hereafter known as ‘PDF<sub>C</sub>’ and ‘PDF<sub>N</sub>’, respectively). PDF<sub>C</sub> was derived from EF values for TOC, DOC,  
242 POC, FA, TEP and CHO. The remaining target compounds derive PDF<sub>N</sub>. In their linear version (Figure 1(a)), both PDFs

243 demonstrate positive skewness (i.e., right-skewness) with the two  $x_m$  values being 1.10 and 1.15, respectively. Nevertheless,  
 244 the peak probability density of the PDF<sub>C</sub> (i.e., the height of the PDF = ~1) is more than twice that of the PDF<sub>N</sub> (~0.4).  $\tilde{x}$  of  
 245 the two PDFs vary substantially, with PDF<sub>C</sub> and PDF<sub>N</sub> yielding values of 1.25 and 1.75, respectively. The values for  $\bar{x}_a$  (2.60  
 246 and 5.39, respectively) and  $\bar{x}_g$  (1.51 and 2.34, respectively) further reflect this divergence. In contrast, their log-transformed  
 247 versions (Fig. 1(b)) approximate normal distributions, with PDF<sub>C</sub> estimating the (exponentials of)  $x_m = 1.11$ ;  $\tilde{x} = 1.24$  and  $\bar{x}_a$   
 248 = 1.51. The PDF<sub>N</sub> yields corresponding values of 1.33, 1.76 and 2.34. Their peak probability densities also reflect that the  
 249 PDF<sub>C</sub> (~2.6) is twice as high as that of PDF<sub>N</sub> (~1.5).



250 **Figure 1: PDFs of the EF values for carbon-enriched (blue) and nitrogen-enriched (orange) compounds.** PDFs of the (a) untransformed  
 251 (i.e., linear) and (b) log-transformed EF values. The solid black line indicates the KDEs derived from original data while the dashed black  
 252 line represents the ensemble mean of bootstrapped KDEs. Central tendency metrics (mode [ $x_m$ ], median [ $\tilde{x}$ ], arithmetic mean [ $\bar{x}_a$ ], geometric  
 253 mean [ $\bar{x}_g$ ]) given in panel (b) are the exponentials of the corresponding estimates on the log scale.

254 We also compared EF-based PDFs (Figure 2) for dissolved (PDF<sub>D</sub>, in purple) and particulate (PDF<sub>P</sub>, in green) OM where we  
 255 refer to a filter size of 0.22  $\mu\text{m}$  (Gao *et al.*, 2012). At a linear scale (Figure 2(a)), the PDFs are again right-skewed for the two  
 256 clusters, with characteristics: (1) 1.20 (both PDF<sub>D</sub> and PDF<sub>P</sub>) for  $x_m$ ; (2) 1.50 and 2.25 for  $\tilde{x}$ ; (3) 4.39 and 4.41 for  $\bar{x}_a$  and, (4)  
 257 1.91 and 2.53 for  $\bar{x}_g$ , respectively. The peak probability density of the PDF<sub>D</sub> (~0.6) exceeds that of the PDF<sub>P</sub> (~0.3) by nearly  
 258 a factor of two. The log PDF<sub>D</sub> and PDF<sub>P</sub> (Fig. 2(b)) approximate normal distributions alongside the following exponentiated  
 259 central values, respectively: (1)  $x_m = 1.20$  and 1.88; (2)  $\tilde{x} = 1.48$  and 2.29; (3)  $\bar{x}_a = 1.91$  and 2.53. Their peak probabilities  
 260 compare between ~1.8 (for PDF<sub>D</sub>) and ~1.0 (for PDF<sub>P</sub>).

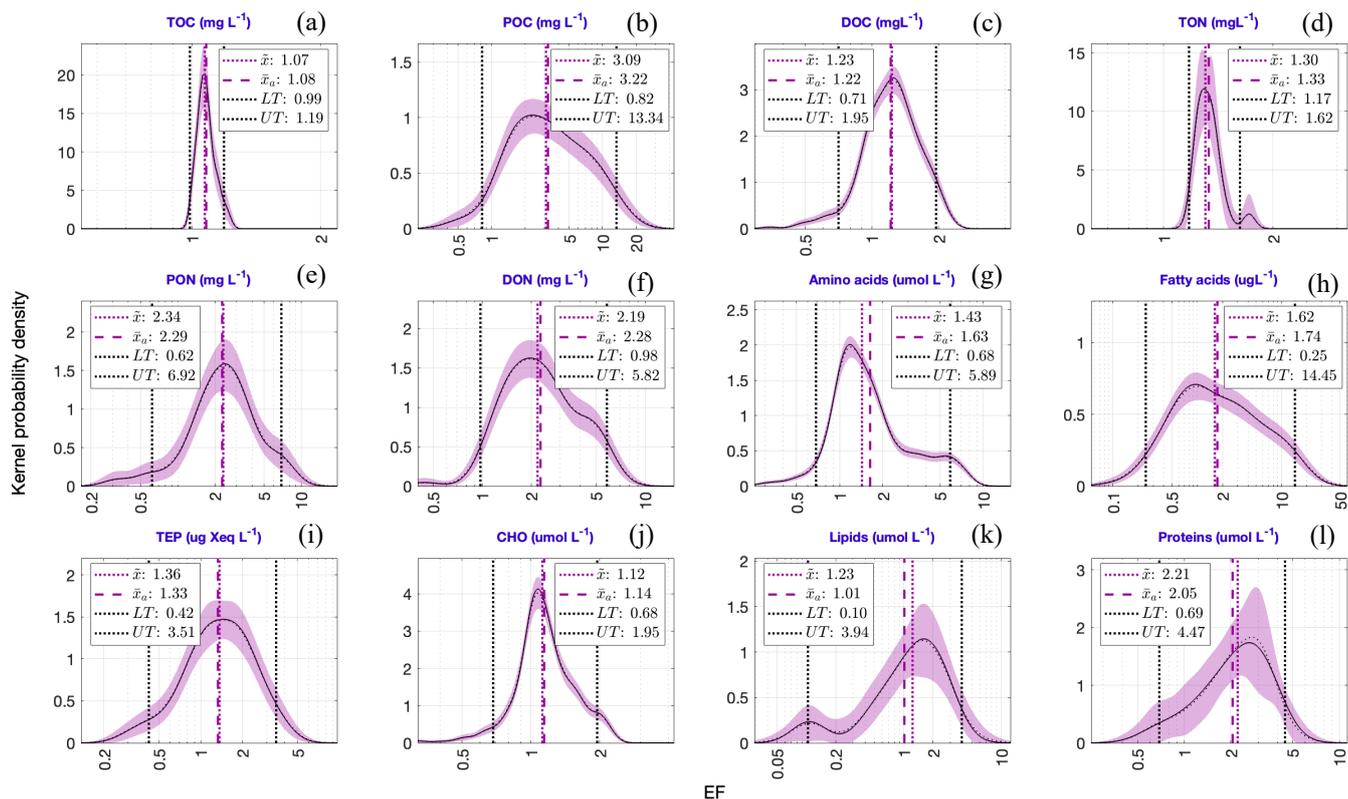


261 **Figure 2: PDFs of the EF values for dissolved (purple) and particulate (green) compounds.** PDFs of (a) linear and (b) log EF values.  
 262 See Fig. 1 caption for details on KDEs and central tendency metrics.

263 Figure 3 displays the PDFs of the EF values for the target compounds. All the distributions exhibit nearly log-normal  
 264 characteristics, nevertheless they vary in their degrees of spread. Here, only  $\tilde{x}$  and  $\bar{x}_a$  values estimate the central tendency of  
 265 each distribution (the rationale for this approach is discussed in section 4.2). The values of  $\tilde{x}$  (dotted pink line) and  $\bar{x}_a$  (dashed  
 266 pink line) are closely aligned in magnitude. According to these derived estimates, median and geometric mean EFs are largest  
 267 for POC (Fig. 3(b):  $\tilde{x} = 3.09$ ;  $\bar{x}_a = 3.22$ ) across all the target compounds, with PON (Fig. 3(e)) and DON (Fig. 3(f)) following  
 268 closely, each exhibiting  $\tilde{x}$  and  $\bar{x}_a$  values  $> 2$ . Although proteins (Fig. 3(l)) also show higher central tendency estimates, it  
 269 should be noted that they have the smallest sample size (= 16), followed by lipids (sample size = 20). Therefore, the results of  
 270 these two compounds should be interpreted with caution due to their lower statistical robustness. A comparison of threshold  
 271 metrics (i.e., LT and UT; see section 2.2.2) reveals that the EF distributions for FA (Fig. 3(h)) and POC (Fig. 3(b)), exhibit the  
 272 highest UT values (14.5 and 13.3, respectively) along with the greatest distributional variability. TOC (Fig. 3(a)) and TON  
 273 (Fig. 3(d)) show the least variability among all target compounds. While some compounds exhibit well-defined unimodal EF  
 274 distributions (e.g., POC, PON), few others (e.g., TON, AA) display polymodal patterns.

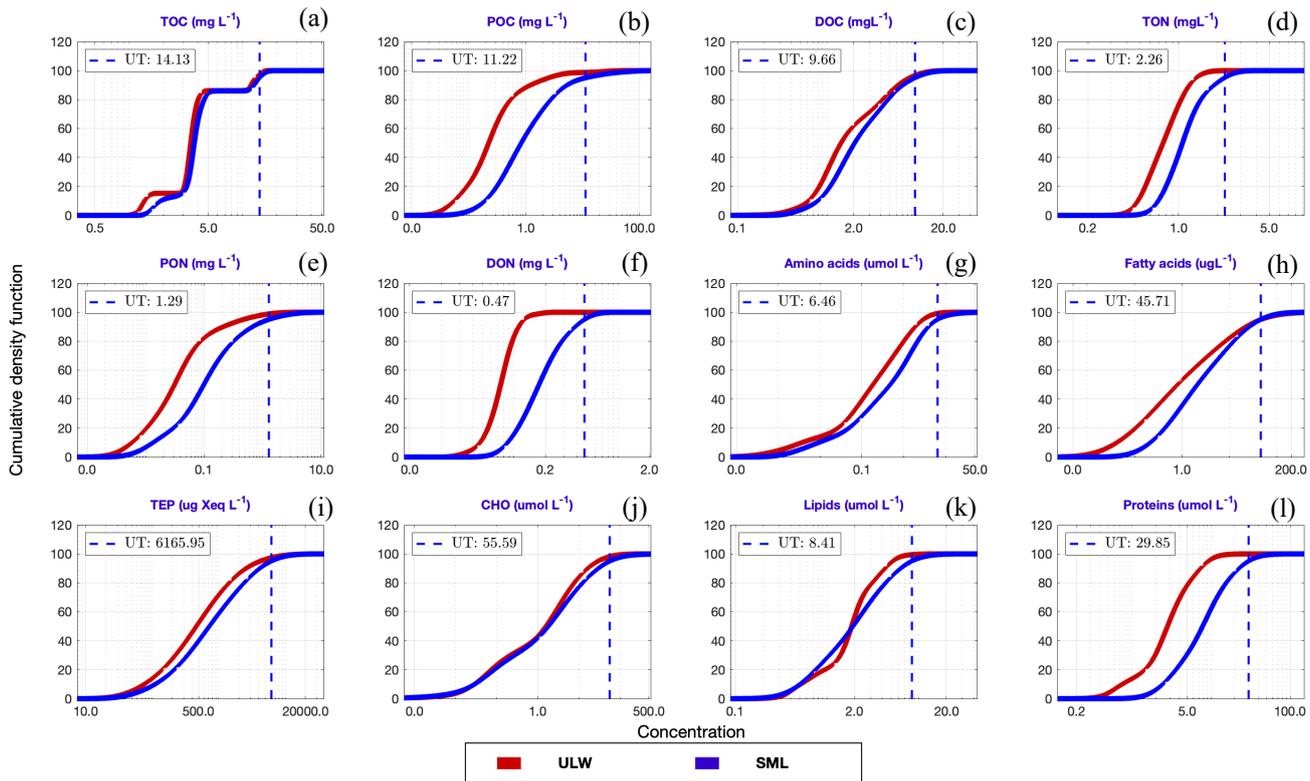
### 275 3.2 Comparing SML and ULW concentrations

276 Figure 4 presents the CDFs of the ULW (in red) and SML (in blue) concentrations for the target compounds. A CDF exhibits  
 277 how probability accumulates across a range of values (in the current context,  $[C]_{\text{SML}}$  and  $[C]_{\text{ULW}}$  data). All CDFs (both  $[C]_{\text{ULW}}$   
 278 and  $[C]_{\text{SML}}$ ) exhibit a characteristic sigmoidal shape: a slow initial rise (i.e., lag phase), followed by a steep rise (i.e.,  
 279 exponential phase), eventually reaching a plateau (i.e., stationary phase). CDFs for TOC display two distinctive plateaus  
 280 indicating bimodal concentration distributions for both SML and ULW.



281 **Figure 3: PDFs of the EF values for the twelve target compounds.** The lower and upper thresholds of each distribution (dashed black  
 282 lines) are defined by 5<sup>th</sup> and 95<sup>th</sup> percentiles of each PDF (see section 2.2.2). The values of these thresholds, along with the central tendency  
 283 metrics given in each panel, are the exponentials of the corresponding estimates in the log space.

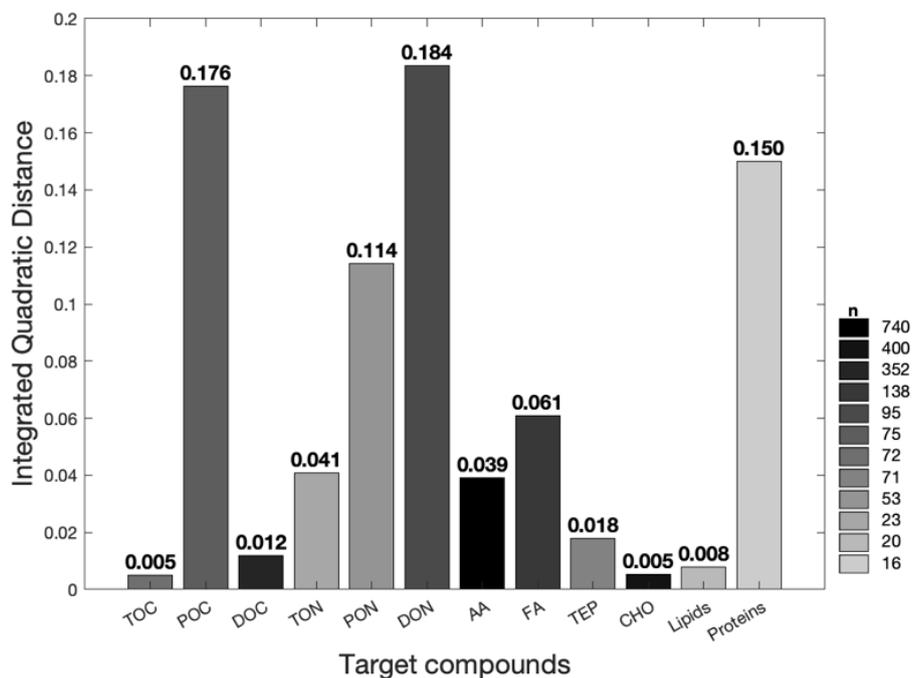
284 Additionally, despite the homogeneity in the general shape and trend of these CDFs, their corresponding IQDs (Fig. 5) reveal  
 285 that the magnitudes of the offsets between  $[C]_{ULW}$  and  $[C]_{SML}$  distributions vary substantially across the target compounds.  
 286 Lower IQD values indicate greater similarities between the CDFs, while higher values document clear distinguishability and  
 287 thus also document a more robust enrichment signal. The lowest IQD is reported for the CDFs of TOC and CHO (0.005) while  
 288 that of DON yield the highest in value (0.184). In addition, lower CDFs (i.e.  $IQD < 0.05$ ) are observed for lipids (0.008), DOC  
 289 (0.012), TEP (0.018), AA (0.039) and TON (0.041), whereas POC (0.18) and proteins (0.15) exhibit a greater divergence (i.e.  
 290  $IQD > 0.15$ ) between  $[C]_{ULW}$  and  $[C]_{SML}$ . The color intensity of each bar reflects the sample size ( $n$ ) of each target compound  
 291 (i.e., smaller the sample, lighter the color).



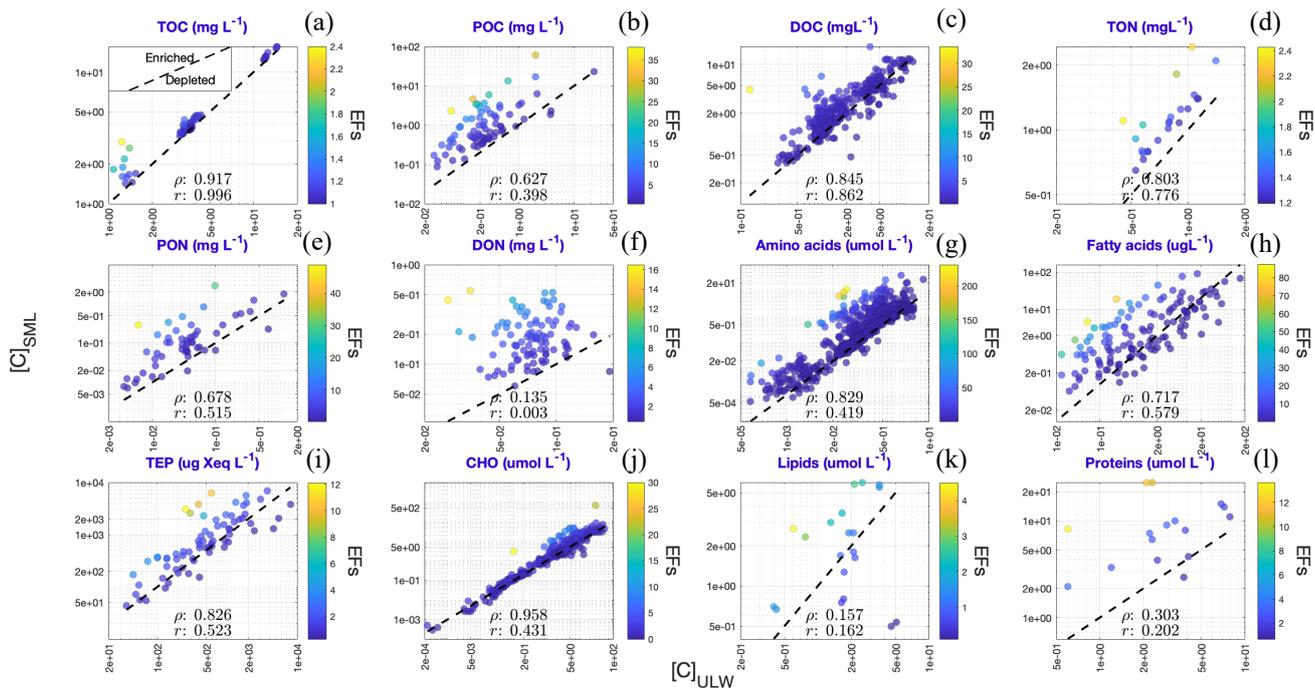
292 **Figure 4: CDFs of the ULW (red) and SML (blue) concentrations for the target compounds.** The upper thresholds for  $[C]_{\text{SML}}$  (UT;  
 293 given by blue dashed lines) are defined by 95<sup>th</sup> percentiles of the corresponding CDF. The values of these thresholds are the exponentials of  
 294 the corresponding estimates in the log space. Their corresponding IQDs are given in Figure 5.

295 Correlations between  $[C]_{\text{ULW}}$  and  $[C]_{\text{SML}}$  of the target compounds were statistically estimated using liner correlation, as  
 296 presented in Figure 6. The coefficients ‘ $\rho$ ’ and ‘ $r$ ’ stand for the correlation values derived from non-parametric Spearman’s  
 297 and parametric Pearson correlation tests, respectively. For all target compounds, except for DON, lipids and proteins, we found  
 298 strong correlations between their SML and ULW concentrations ( $\rho$  and  $r > 0.5$ ) with robust positive relationships. Individual  
 299 datapoints for TOC, DOC and CHO (Figs. 6(a), (c) and (j)) closely fall on the 1:1 reference line where  $[C]_{\text{SML}} = [C]_{\text{ULW}}$  (dashed  
 300 black line). In contrast, those for POC, TON, AA and FA are notably shifted towards the y-axis, suggesting higher  $[C]_{\text{SML}}$   
 301 values relative to  $[C]_{\text{ULW}}$  that corresponds to potentially enriched (depleted) SML (ULW) concentrations against ULW (SML)  
 302 concentrations (see inset plot in Fig. 6(a)). Although TEP shows a slight enrichment in the SML, the effect is not particularly  
 303 pronounced (Figure 6(i)). In addition, all the datapoints (regardless of whether they display copulation or not) were further  
 304 color-coded according to their respective EFs. The results reveal an overall consistency in EFs across concentration ranges

305 irrespective of their magnitudes. For example, in Fig. 6(c), EF values remain below 5, both when  $[C]_{ULW}$  and  $[C]_{SML}$  are  $< 0.5$   
306  $\text{mg L}^{-1}$  and  $> 5 \text{ mg L}^{-1}$ . This pattern holds across nearly all the target compounds.



307 **Figure 5: IQD values quantifying the divergence between ULW and SML concentrations for each target compound.** The IQD  
308 represents the squared difference between ULW- and SML-based CDFs shown in Figure 4. Higher IQD indicates greater divergence between  
309 the two distributions and vice versa. Bar color intensity corresponds with the sample sizes.



310 **Figure 6: Linear correlation between  $[C]_{ULW}$  and  $[C]_{SML}$  for the target compounds.** The datapoints are color-coded based on their  
 311 corresponding EFs. Dashed black line indicates 1:1 line when  $[C]_{ULW}$  (x-axis) =  $[C]_{SML}$  (y-axis). Inset plot in panel (a) exhibits the relevant  
 312 implications of the figure: Correlations above the 1:1 line corresponds to selective SML enrichment and vice versa. The values of 'ρ' and  
 313 'r' give Spearman's and Pearson correlation coefficients, respectively.

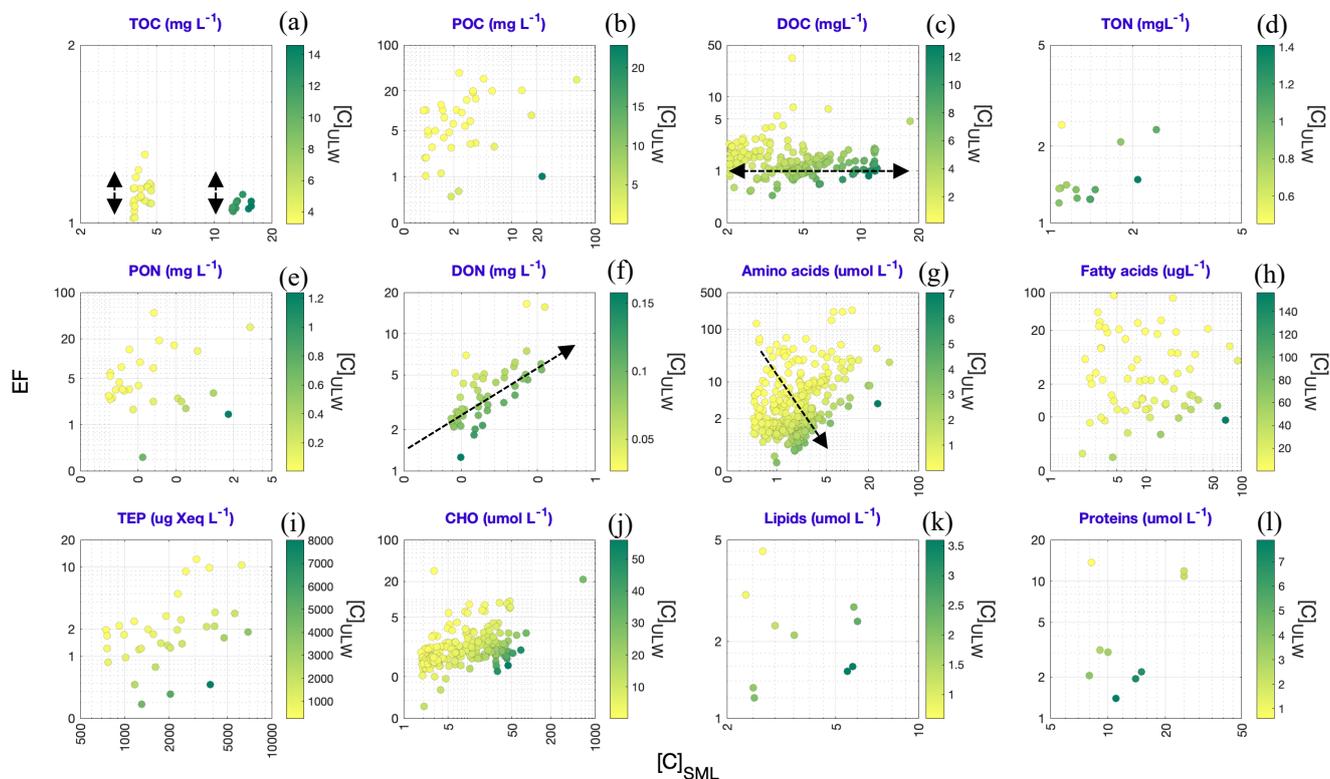
### 314 3.3 Investigating concentration-dependent enrichment dynamics

315 Informed by the observations drawn from Fig. 6, Fig. 7 presents a more detailed investigation into the interrelationships among  
 316  $[C]_{SML}$ ,  $[C]_{ULW}$  and EFs in the environment. The analysis is restricted to  $[C]_{SML}$  values (x-axis) that exceed the  $\tilde{x}$  (i.e., median)  
 317 of their respective distributions (median is the most stable central tendency metric of a distribution. Discussed further in section  
 318 4.2). These elevated  $[C]_{SML}$  are compared against the corresponding  $[C]_{ULW}$  (color scale) and EF (y-axis) values. The results  
 319 reveal following covariation trends:

- 320 (i) TOC reports a generally low range of EF values comparable at both low and high concentrations of SML and  
 321 ULW (Fig. 7(a))
- 322 (ii) DOC displays relatively consistent EF values regardless the magnitudes of  $[C]_{SML}$  and  $[C]_{ULW}$  (Fig. 7(c)), but  
 323 also slightly points towards higher EF values in association with low  $[C]_{ULW}$
- 324 (iii) DON presents an ascending EF gradient, positively correlated with  $[C]_{SML}$  (Fig. 7(f)), revealing more enrichment  
 325 to be well reflected in the concentrations found in the SML

326 (iv) AA shows a similar correlation dependence to that of DON, but also reveals a much clearer trend toward higher  
327 EF values to be found at lower  $[C]_{ULW}$  concentrations (Fig. 7(g))

328 FA (Fig. 7(h)), despite their larger sample sizes, exhibit no clear trend in the  $[C]_{SML} - [C]_{ULW} - EF$  triad.



329 **Figure 7: Interdependent relationship of  $[C]_{SML}$  values with the corresponding  $[C]_{ULW}$  and EF values.** The analysis is restricted to  
330  $[C]_{SML}$  values that exceed the corresponding  $\bar{x}$  values. The x-axes give the observed  $[C]_{SML}$  values against their corresponding EF values on  
331 y-axes. Datapoints are color-coded based on corresponding  $[C]_{ULW}$  values. The black arrows indicate identified enrichment patterns.

#### 332 4. Discussion

333 A major strength of the SML-OM dataset is its broad coverage of OM concentrations, whereas individual studies are typically  
334 restricted to a narrow range of similar ULW conditions. By employing a meta-analytical approach, our study presents the first  
335 comprehensive overview of the enrichment dynamics in the SML, based on existing literature. Meta-analytical studies offer a  
336 rigorous framework to synthesize evidence across diverse datasets thereby improving the reliability of scientific conclusions.  
337 By statistically integrating outcomes from independent investigations, meta-analyses increase overall analytical power, reduce

338 the influence of small-sample variability and, uncover true environmental signals from study-specific biases (i.e., sampling  
339 strategies, analytical techniques, spatial scales and seasonality). This not only enhances the generalizability of findings but  
340 also exposes gaps and inconsistencies in the existing literature, guiding the development of more robust future studies.

341 In this context, conducting a quantitative assessment of how the reference studies (i.e., those on which the SML-OM is based)  
342 are distributed across key domains of SML research provides insights into the most frequently studied aspects (Supplementary  
343 Figure S1), thereby explicitly quantifying metadata coverage; research on the SML has increased from about 15 publications  
344 per year in the early 2000s to approximately 50 per year by 2016 (Engel *et al.*, 2017). However, our work highlights the  
345 potential understudied areas in SML research that call for more in-depth analysis. For instance, majority of the reference studies  
346 has been conducted in oceanic and coastal regions (~76% of data) and predominantly during warmer months (~77% of data)  
347 with a significant mismatch observed for data collected under low and high wind regimes (~81% vs. 19%, respectively). In  
348 light of these research gaps, the following sections interpret the main findings revealed by our analysis and discuss their  
349 implications for understanding SML enrichment.

## 350 4.1 Overarching trends in SML enrichment

### 351 4.1.1 Generalized enrichment patterns

352 Comparison of KDE-derived PDFs for the EF values of (1) carbon-enriched vs. nitrogen-enriched organic compounds (Fig. 1)  
353 and (2) dissolved vs. particulate organic compounds (Fig. 2) yield the following key implications:

- 354 (1) All the estimated original (i.e., linear scale) PDFs (Figs. 1(a) and 2(a)) display higher probability densities for lower  
355 EF values and extended tails towards higher EF values (i.e., right-skewness), suggesting that under natural conditions,  
356 modest SML enrichment is far more common in general, while extreme enrichment events are rare.
- 357 (2) Nevertheless, variation in the peak probability densities among the PDFs indicate that extreme SML enrichment  
358 events are relatively more frequent in nitrogen-enriched compounds (Fig. 1: orange PDF) and particulate forms (Fig.  
359 2: green PDF), compared to carbon-enriched compounds (Fig. 1: blue PDF) and dissolved forms (Fig. 2: purple PDF)
- 360 (3) Nitrogen-enriched compounds and particulate forms exhibit a broader EF variability (i.e., higher mode, median, mean  
361 values) compared to carbon-enriched compounds and dissolved forms with a relatively more consistent spread (i.e.,  
362 lower central tendency metrics)

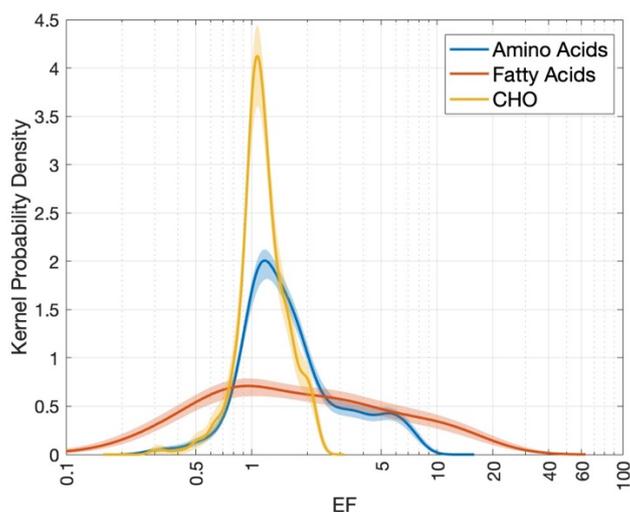
363 These differences in peaks and central tendency metrics persist in log-transformed PDFs as well (Figs. 1(b) and 2(b)). This  
364 validates that these variations are not caused by statistical artifacts but reflect real, natural variability in enrichment behavior.  
365 Overall, these findings from our meta-analysis indicate that the OM accumulation in the SML is more effective for (1) nitrogen-  
366 enriched than for carbon-enriched compounds and (2) particulate than for dissolved forms. *These enrichment patterns likely*

367 reflect the combined influence of biological, chemical and physical mechanisms acting on the SML: Frka *et al.* (2012) outlined  
368 the complex, multicomponent behaviour of the SML, suggesting competitive adsorption of more insoluble surfactants (e.g.,  
369 lipid-like material) during highly productive seasons. Selective enrichment of certain polar amino acids (e.g., arginine and  
370 glutamic acid) relative to others has been hypothesized (Barthelmeß and Engel, 2022). However, these mechanisms are not  
371 yet fully resolved. The molecular structure of N-enriched compounds also promotes aggregation into colloids or microgel  
372 particles (Dietz *et al.*, 1976), a process further enhanced by bubble scavenging and low-turbulence trapping (Mopper *et al.*,  
373 1995). In contrast, carbon-rich compounds such as polysaccharides are generally more soluble and tend to remain largely  
374 distributed in the bulk water, resulting in comparatively lower surface activity and enrichment at the interface (Ćosović and  
375 Vojvodić, 1989; Laß *et al.*, 2013). Similarly, particulate OM exists as discrete, larger units, that can be trapped at the interface  
376 due to surface tension and by bubble-mediated processes (Robinson *et al.*, 2019). However, some particles, such as TEP  
377 ballasted with mineral dust or phytoplankton shells, may sink rather than rise, highlighting the complex balance of forces  
378 controlling surface accumulation (Mari *et al.*, 2017). While bubble scavenging represents a process, which can lead to the  
379 aggregation of dissolved components (for example, at the rear of rising bubbles; Dukhin *et al.*, 2015; Zhou *et al.*, 1998), a  
380 large fraction of dissolved OM, potentially exhibiting reduced overall surface activity, passes through these processes without  
381 accumulating at the air–water interface. Transient enrichment of dissolved OM can nonetheless occur. Overall, surface-  
382 associated processes such as bubble scavenging and aggregation can enhance the enrichment of particulate compounds in the  
383 SML relative to dissolved forms. .

384 Our findings contradict some earlier works, including Baastrup-Spohr and Staehr (2009), Liss and Duce (1997), Yang (1999),  
385 who suggest that the SML is similarly enriched for both particulate and dissolved organic (and inorganic) compounds, but  
386 align with other studies that report opposing results: Dietz *et al.* (1976) provide evidence for enhanced accumulations of  
387 particulate matter in the SML through particle aggregation at the surface. The work further links high abundances of living  
388 bacteria in the near-surface to higher availability of POC in the SML. Studies of Carlucci *et al.* (1985), Henrichs and Williams,  
389 (1985), Kuznetsova *et al.* (2004), Kuznetsova and Lee (2002) and Reinthaler *et al.* (2008) report that POC and PON tend to  
390 be more enriched in the SML than DOC. Engel *et al.* (2017) state that the SML has been shown to be enriched in particulate  
391 organic matter, particularly in proteinaceous compounds. Together, these findings and our meta-analytical synthesis, indicate  
392 that nitrogen-enriched compounds may interfere more critically with interfacial properties. Whether incorporating these  
393 nitrogen-based metrics can improve the precision of traditional carbon-only parameterizations of gas exchange suppression  
394 (e.g., Barthelmeß *et al.*, 2021; Li *et al.*, 2024) remains an important question for future work.

395 Although nitrogen-enriched compounds seem to exert a strong influence on the SML's interfacial properties at the bulk scale,  
396 a compound-specific comparison of three biosurfactants data – AA (Fig. 4(g)), FA (Fig. 4(h)) and CHO (Fig. 4(j)) – reveals a  
397 more complex picture (Figure 8): AA, despite their high nitrogen content, does not exhibit the highest enrichment. Instead, its  
398 EF values (0.3 - 10) fall between those of CHO (lower end: 0.3 – 2.5) and FA (upper end: 0.1 - 60). Our results are consistent

399 with earlier reports on their natural EF ranges: Polysaccharides can be enriched in the SML up to three-fold compared to the  
400 ULW (Williams *et al.*, 1986; Wurl and Holmes, 2008). Enrichment of AA can vary between 0.3 to 201, depending on their  
401 species-specificity (Cunliffe *et al.*, 2013). This EF hierarchy likely reflects the intermediate surface activity of these three  
402 compounds: (Barthelmeß and Engel, 2022), referring to Ćosović and Vojvodić (1998) state that “*Lipid-like surfactants exhibit*  
403 *stronger surface activity, while protein-like, followed by carbohydrate-like, surfactants decrease in activity*”. FA dominate  
404 competitive adsorption due to their strong amphiphilic character (i.e., a long hydrophobic tail and small polar head), allowing  
405 them to readily form stable monolayers at the interface. In contrast, proteins are moderately surface-active, while highly soluble  
406 CHO largely remain in the bulk water (Ćosović and Vojvodić, 1998; Laß *et al.*, 2013; Laß and Friedrichs, 2011). Consequently,  
407 these patterns suggest that compound-specific enrichment in the SML is driven more by the surfactant properties of individual  
408 compounds than their elemental composition alone. While this hierarchy holds for single-component systems, interactions in  
409 natural SML mixtures are complex, and the structural properties of the nanolayer are influenced by both rare insoluble lipid-  
410 like and abundant soluble carbohydrate-like material. Nevertheless, we do not extend this analysis further as biosurfactants  
411 measurements (1) primarily quantify surface activity rather than enrichment and (2) are subjected to methodological  
412 inconsistencies that limit cross-study comparability. We further acknowledge that the mechanistic links between surfactants,  
413 their behavior and associated ecosystem processes remain incompletely understood.



414 **Figure 8: Comparison of PDFs of the EF values for three bio-surfactants: Amino acids (blue), fatty acids (orange) and carbohydrates**  
415 **(yellow).** The figure synthesizes the KDEs given in Figs. 3(g), (h) and (j).

416 Overall, these overarching trends of SML enrichment underscore the importance of resolving compound-specific accumulation  
417 in the SML, while distinguishing between selective and non-selective enrichment. Cumulative probability comparison results

418 for the  $[C]_{ULW}$  and  $[C]_{SML}$  (Figs. 4 and 5) and their corresponding linear correlations (Fig. 6) provide a meta-analytical  
419 perspective on how compounds are distributed and accumulated between the two compartments. Here, results concerning  
420 lipids and proteins are excluded due to apparent randomness in their distributions, potentially caused by smaller sample sizes.

#### 421 4.1.2 Compound-specific enrichment patterns

422 Significant correlations between  $[C]_{ULW}$  and  $[C]_{SML}$  of nearly all the target compounds ( $\rho$  and  $r > 0.5$ ) are consistent with the  
423 overall understanding that the SML's composition is linked to the availability of material in the underlying sub-surface waters  
424 (Chen *et al.*, 2016; Joux *et al.*, 2006; Kuznetsova *et al.*, 2004). Contrary to this general pattern, Kuznetsova *et al.* (2004)  
425 suggest that certain OM fractions in the SML and ULW may show lack of correlation, potentially due to constraints such as  
426 varying mineralization rates between the two layers and surface adsorption processes. Consistent with this view, linear  
427 correlation results for DON indicate such decoupling (Fig. 6(f)), though the underlying causes remain unexplored in this study.  
428 Early works also suggested that the variations in the SML concentrations are typically larger than those in the ULW (Reinthal  
429 *et al.*, 2008). In agreement, CDFs of the  $[C]_{ULW}$  and  $[C]_{SML}$  demonstrate faster probability accumulation for ULW than SML  
430 (Fig. 4), implying generally smaller magnitudes and lower variability in ULW concentrations compared to SML  
431 concentrations. Conversely, Carlson (1983) argues that, in certain occasions, OM variability in the SML and ULW may not  
432 significantly differ across temporal and spatial scales. The CDFs for TOC, DOC, TEP and CHO which exhibit the lowest IQD  
433 values (Fig. 5), support this but is contradicted by those of the other compounds, with higher IQD values (indicating substantial  
434 differences between the two concentrations).

435 Works of Hunter and Liss (1977) and Kurata *et al.* (2016) discuss the selective enrichment of surfactants in the SML, mainly  
436 driven by microbial processes. Hydrophobic compounds tend to show more affinity to the surface compared to hydrophilic  
437 substances (Marty and Saliot, 1976). In agreement, our linear correlation results reveal preferential accumulation of **the**  
438 **biosurfactants**, AA (Fig. 6(g)) and FA (Fig. 6(h)), in the SML. Linear correlation results shown in Fig. 6(i) provide evidence  
439 to the view that TEP is generally enriched in the SML compared to the ULW (Cunliffe and Murrell, 2009; Cunliffe *et al.*,  
440 2009; Wurl and Holmes, 2008), although this enrichment is not strongly pronounced in our dataset. Additionally, the nearly  
441 overlapping CDFs for TEP in SML and ULW (Fig. 4(i)) along with its low IQD value (= 0.081; Fig. 5) **indicate a surprisingly**  
442 **weak enrichment, contrary to expectations. Nevertheless,** concentration trend of TEP observed in our data closely aligns with  
443 that of CHO (Figs. 4(j) and 5), supporting the prevailing hypothesis that TEP is formed through coagulation of dissolved  
444 polysaccharides (Passow, 2000). Thornton *et al.* (2016) observe that TEP and dissolved polysaccharides do not always exhibit  
445 significant enrichment in the SML as anticipated.

446 POC and PON correlation patterns (Figs. 6(b) and 6(e), respectively) where  $[C]_{SML}$  significantly exceed  $[C]_{ULW}$ , and that of  
447 DOC (Fig. 6(c)) where  $[C]_{SML}$  is nearly equal to  $[C]_{ULW}$ , provide strong meta-analytical evidence to earlier works that discuss  
448 the selective enrichment of POC and PON in the SML over DOC (e.g., Carlucci *et al.*, 1992; Henrichs and Williams, 1985;

449 Kuznetsova and Lee, 2002; Kuznetsova *et al.*, 2004 and Reinthaler *et al.*, 2008). Carlson (1983) suggests that the distribution  
450 of some organic fractions between the SML and the ULW may be governed by specific partitioning processes. For instance,  
451 while Chen *et al.* (2016) point out the significant role of the ULW in DOC and CHO accumulation in the SML, Dietz *et al.*  
452 (1976) observe fairly consistent abundances for these compounds between the two layers. Our experiments also show strong  
453 1:1 correlation for DOC (Fig. 6(c)) and CHO (Fig. 6(j)), suggesting an absence of preferential affinity towards the SML (unlike  
454 surfactants), which further indicates that their enrichment is predominantly controlled by the ULW. Although CHO, AA and  
455 FA are identified to be the key constituents of the organic carbon pool (Hedges *et al.*, 1994), our correlation results reveal that  
456 their partitioning between the SML and the ULW and, their eventual enrichment patterns, may not be consistent (Figure 6), as  
457 also suggested by Fig. 8.

#### 458 4.1.3 Influencing factors and current uncertainties

459 Baier *et al.* (1974), Hunter and Liss (1981) and MacIntyre (1974) argue that the compositional diversity of the SML prevents  
460 single compounds from fully representing the dissolved OM class, which further emphasizes the importance of assessing  
461 compound-specific accumulation in the SML. Such investigations could shed light on selective and non-selective enrichment  
462 dynamics of OM. An analysis of EF-based PDFs for various AA fractions (Figure 9(a)) – Total AA (TAA), Dissolved Free  
463 AA (DFAA), Dissolved Combined AA (DCAA) and Particulate AA (PAA) – revealed notable heterogeneity within this  
464 compound class, reflecting the chemical diversity and complexity of OM enrichment in the SML: *Relatively lower enrichment*  
465 *in DFAA may indicate its limited accumulation in the SML, potentially due to its high solubility and rapid turnover (Jørgensen*  
466 *et al., 1993). In contrast, DCAA, which comprises combined amino acids, might exhibit stronger surface activity and a greater*  
467 *tendency to form aggregates, leading to higher enrichment across a broader range. The bimodal EF distribution observed for*  
468 *PAA could reflect differences in particle composition, size and hydrophobicity, whereby denser particles sink rapidly while*  
469 *buoyant, organic-rich particles preferentially accumulate at the surface. TAA, which integrates all these molecular states, may*  
470 *dampen these extremes and yields more moderate enrichment. These interpretations remain as hypotheses, as very little is*  
471 *known about the behavior of AA embedded in highly complex structures in natural SML.*

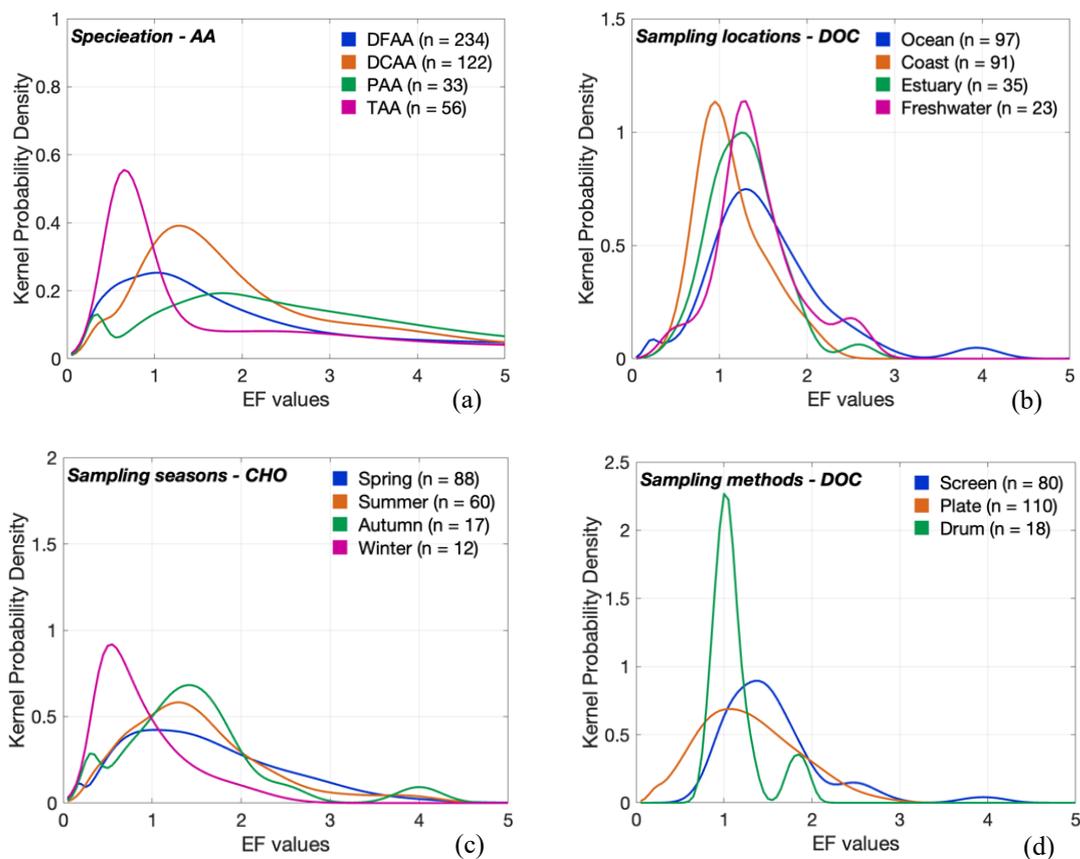
472 Additionally, consistent with previous studies that investigated the influence on environmental drivers on the enrichment  
473 dynamics in the SML (e.g., Asher, 1997; Barthelmeß *et al.*, 2021; Knulst *et al.*, 1997; Kuznetsova *et al.*, 2004; Liu and Dickhut,  
474 1998; Obernosterer *et al.*, 2008; Reinthaler *et al.*, 2008; Tsai and Liu, 2003), our analysis demonstrates that factors such as  
475 sampling location (for DOC), sampling season (for CHO) and sampling method (for DOC) (Figs. 9(b) – (d)) play key roles in  
476 modulating the enrichment variability of the OM. It should be noted that these specific target compounds are chosen as  
477 representative examples because they span all subcategories of secondary data considered in the study (see Supplementary  
478 Table S1), and therefore enable a more robust comparison among different settings.

479 Comparison of DOC enrichment across four sampling locations (Fig. 9(b)) reveals relatively reduced coastal enrichment (as  
480 also seen by Carlson, 1983). This pattern may result from enhanced mixing and shorter surface residence times due to stronger  
481 wave action, tidal influence, or nearshore turbulence. Alternatively, the limited two-dimensional space at the air-sea interface  
482 in combination with a highly saturated underlying water layer could also contribute to reduced enrichment. At the same time,  
483 broader EF variability displayed by oceanic sites likely stems from the greater heterogeneity of open-ocean conditions  
484 including varying biological productivity and OM sources (Carlson, 1983). In contrast, estuarine and freshwater systems,  
485 which often have more constrained physical regimes and relatively consistent OM inputs, tend to exhibit narrower EF ranges  
486 (Hillbricht-Ilkowska and Kostrzevska-Szlakowska, 2004). Observations by Barthelmeß *et al.* (2021) and Mustaffa *et al.*  
487 (2018) suggest that changes in EFs in these systems are driven more by variations in the ULW than by the SML itself, consistent  
488 with the idea that the surface layer in high-OM regimes is already saturated and thus less responsive to additional inputs.  
489 Moreover, bimodal CDFs of TOC for the SML and ULW concentrations (Fig. 4(a)) along with the distinct separation of three  
490 data clusters in their correlation patterns (Fig. 6(a)), further highlight the significant role of spatial factors in shaping SML  
491 composition. The origins of TOC data used in this study illustrates this variability: Data from (1) a heavily polluted urban lake  
492 (concentration range: 12 – 16 mg L<sup>-1</sup>; Baastrup-Spohr and Staehr, 2009), (2) a forested lake (concentration range: 3 – 5 mg L<sup>-1</sup>;  
493 Baastrup-Spohr and Staehr, 2009), (3) the Arctic Ocean (concentration range: 1 – 3 mg L<sup>-1</sup>; Gao *et al.*, 2012) and (4) an  
494 upwelling filament (concentration range: 3 – 4 mg L<sup>-1</sup>; Barthelmeß *et al.*, 2021).

495 Seasonal comparison of the EF values for CHO (Fig. 9(c)) likely reflects temporal differences in biological drivers (Gašparović  
496 and Čosović, 2001; 2003) in the SML (i.e., depletion in winter, while the other seasons show relatively higher and consistent  
497 enrichment with broader variability). Additionally, stronger wind conditions typical of winter may disturb the SML and reduce  
498 particle residence time, counteracting surface accumulation (Sun *et al.*, 2018). However, the influence on wind speed on SML  
499 enrichment remains ambiguous; our comparison of EF values under calm (< 6.6 ms<sup>-1</sup>; Reinthaler *et al.*, 2008) and rough (> 6.6  
500 ms<sup>-1</sup>) wind conditions yield inconclusive results (Supplementary Figure S2) with wind speed appearing to have little/no effect  
501 on the SML enrichment (e.g., Baastrup-Spohr and Staehr, 2009 and Sabbaghzadeh *et al.*, 2017) or with enrichment persisting  
502 even under rough sea conditions (e.g., Kuznetsova *et al.*, 2004; Reinthaler *et al.*, 2008), opposing the general understanding  
503 that turbulent conditions may reduce the concentration in the SML (e.g., Carlson, 1983).

504 Nevertheless, it is important to note that imbalanced sampling efforts among these categories, (Supplementary Figure S1),  
505 specially with regards to wind speed (Supplementary Figure S2; see the sample sizes), may compromise the robustness and  
506 validity of these findings. This is also evident in the comparison of sampling techniques for the EF values of DOC (Fig. 9(d)),  
507 where the drum method, with a sample size of only 17, shows limited variability compared to the screen and plate methods  
508 (sample sizes of 80 and 110, respectively). This likely reflects a bias due to sampling effort rather than a true difference in  
509 enrichment behavior. Collectively, these findings highlight the need for future SML research and SML-based model

510 development to systematically account for wind and sea-state conditions, and to explore how enrichment patterns may vary  
511 under different environmental regimes and methodological settings.



512 **Figure 9: Factor-specific enrichment variability in the SML.** PDFs illustrating varying enrichment patterns for (a) AA across chemical  
513 forms, (b) DOC across sampling locations, (c) CHO across sampling seasons and (d) DOC across sampling methods. 'n' gives the sample  
514 size of each category. Supplementary Table S1 summarizes different sampling locations, sampling seasons and sampling methods observed  
515 for the investigated target compounds.

516 Another major source of uncertainty arises from the variability in sampling depths of the ULW (Supplementary Table S1),  
517 which can affect the comparability of different data across multiple studies that would eventually introduce bias into the  
518 interpretation of overarching trends. Additional biases which are beyond the scope of this study include the potential influence  
519 of diurnal cycles (López-Puertas et al., 2025); OM can be rapidly removed from the SML through photochemical degradation  
520 (Obernosterer et al., 2008) and also be affected by reduced bacterial metabolism due to solar radiation (Dietz et al., 1976).  
521 Therefore, taken together, our meta-data analysis suggests that, investigating SML enrichment without accounting for these

522 influencing factors may mask true enrichment patterns, limiting the ability to derive meaningful insights. In light of these  
523 considerations, our work highlights the need for conducting species-specific and condition-dependent analyses in future SML  
524 research that also focus on subsequent environmental parameters, as also proposed by Pereira et al. (2018).

#### 525 **4.2 Scale-related biases in EF estimates**

526 Accurate data interpretation is essential to gain precise insights and arrive at substantiated conclusions (Isles, 2020; Menge *et*  
527 *al.*, 2018). This is particularly true for meta-analyses involving continuous environmental data where values may vary by  
528 several orders of magnitude (e.g., Vitousek, 2004). In our study, when the PDF<sub>C</sub> vs. PDF<sub>N</sub> (Fig. 1) and PDF<sub>D</sub> vs. PDF<sub>P</sub> (Fig.  
529 2) are evaluated on a linear scale (panels (a)), they exhibit right-skewness, whereas their log-transformed versions approximate  
530 normal distributions (panels (b)). Comparisons between highly skewed distributions raise uncertainties as their offsets are  
531 often dominated by extreme values/outliers. In contrast, when log transformation is applied, the distributions tend to exhibit  
532 more symmetric, normalized patterns which enable direct comparisons in shape and spread across different categories (Zuur  
533 *et al.*, 2007). Therefore the normality assumption for EF is inappropriate and the computation of an arithmetic mean, a  
534 conventional practice adopted in many earlier works (e.g., Gašparović *et al.*, 2007; Gao *et al.*, 2012; Kuznetsova *et al.*, 2005;  
535 Williams *et al.*, 1986; Wurl and Homes, 2008; Wurl *et al.*, 2009), can be misleading, likely providing a biased general picture  
536 of OM enrichment in the SML.

537 The here constructed PDFs given in Figures 1 and 2 reveal that both mode ( $x_m$ ; shown by solid straight lines) and arithmetic  
538 mean ( $\bar{x}_a$ ; shown by dashed straight lines) differ between the two scales: The mode reflects the peak of a distribution and is  
539 sensitive to the shape of its respective density curve. It varies depending on whether a dataset is in ‘skewed’ linear space or  
540 ‘normalized’ log space and becomes ambiguous in polymodal distributions (regardless of the scale: e.g. Fig. 3). As a  
541 consequence, the mode in general provides an unreliable measure of central tendency. While the linear-arithmetic mean, which  
542 is influenced by outliers, result in biases that exaggerate the corresponding central tendency, the log-arithmetic mean prevents  
543 the extreme values from being dominant through balanced averaging and hence provides a reliable estimation of central  
544 tendency. Nevertheless, geometric mean in linear space ( $\bar{x}_g$ ; straight lines with alternating dots and dashes) is a meaningful  
545 measure given that it is equivalent to the exponential of the arithmetic mean in logarithmic space (See Eqs. (B1) and (B2)).  
546 Median ( $\tilde{x}$ ; dotted straight lines), on the other hand, remains relatively consistent across both scales as it is a rank-based  
547 measure of central tendency that is unaffected by the magnitude of outliers. Accordingly, we suggest that future SML  
548 enrichment studies employ a logarithmic scale for data analyses, and adopt either geometric mean and/or median on linear  
549 scale or arithmetic mean and/or median on logarithmic scale for reliable trend analysis.

550 Based on these new insights on scale transformations and central tendency metric considerations, we have redefined the typical  
551 EF values of the studied target compounds and their degrees of spread from a meta-analytical perspective, from the estimated  
552  $\tilde{x}$ ,  $\bar{x}_a$  and thresholds (i.e. UT and LT) of their PDFs (Fig. 3). To re-establish these EF ranges as generally observed estimates

553 under common conditions, ‘the box plot method’ (Tukey, 1977) was applied to the data to detect and remove potentially  
554 extreme EF values that rarely occur in nature. By providing these systematically derived ranges, our analysis offers a robust  
555 and comprehensive reference framework, enabling future SML-based studies to consistently evaluate newly obtained EF  
556 measurements, assess their position relative to typical distributions, and identify deviations that may indicate unusual  
557 environmental conditions or methodological inconsistencies.

### 558 4.3 Role of EF in reflecting SML enrichment

559 While the metric of EF offers a convenient way to assess the accumulation trends in the SML and therefore serves as the basis  
560 for many established insights and inferences in existing SML research (see Introduction), its ability to accurately and robustly  
561 express the ‘true’ enrichment nature of the SML has constantly been a question of interest (e.g., Basstrup-Spohr and Staehr,  
562 2009; Hillbricht-Ilkowska & Kostrzewska-Szlakowska, 2004; Knulst *et al.*, 1997; Kuznetsova *et al.*, 2004; Liss and Duce,  
563 1997; Münster *et al.*, 1998; Södergren, 1987). The EF is a ratio that expresses the ‘relative’ changes in  $[C]_{\text{SML}}$  with respect to  
564  $[C]_{\text{ULW}}$  (Eq. (1)), and hence is sensitive to the variations in either layer. Ideally, to effectively reflect conditions of growing  
565 SML enrichment, EF values should gradually rise in response to increasing  $[C]_{\text{SML}}$  and decreasing  $[C]_{\text{ULW}}$ , which can be visibly  
566 observed for DON (Fig. 7(f)), AA (Fig. 7(g)) and CHO (Fig. 7(j)). Nevertheless, our meta-analysis highlights several  
567 inconsistencies that challenge the relevance of the EF values as indicators of ‘true’ SML enrichment. For instance, on one  
568 hand, similar EF values can be observed for both oligotrophic and eutrophic environments (referring to the EFs of TOC: Fig.  
569 7(a)), which limits the ability to distinguish the differences in their trophic status (i.e., nutrient/productivity characteristic of  
570 the water body), despite them being conspicuous in TOC’s absolute concentration range (bimodal CDFs; Fig. 4(a)). On the  
571 other hand, high (low) EF values may occur under oligotrophic (eutrophic) conditions leading to over- (under-) estimation of  
572 ecological setting (i.e., biological and environmental context under which SML samples were collected; Fig. 7(g)).  
573 Furthermore, symmetrical changes in SML and ULW yield near-constant EF values across a wide range of concentrations  
574 (Fig. 7(c)), which could cause misinterpretations in key ecosystem shifts. We have also observed consistent EF values, even  
575 when SML and ULW concentrations vary over several orders of magnitudes (Figs. 6(g) – (j)), which further raise concerns  
576 over the metric’s robustness. Therefore, although widely used, EF values should be interpreted with caution and, combined  
577 with additional parameters that provide more accurate information about the true enrichment behaviour of the SML.

578 A complementary parameter would be the typical upper limit of a  $[C]_{\text{SML}}$  distribution which may reflect the maximum  
579 concentration capacity of the SML. Such a measure can serve as a robust concentration estimate of such maximum capacity if  
580 approximated from a meta-data derived distribution that includes observations across all diverse environmental conditions.  
581 Table 1 summarizes the upper  $[C]_{\text{SML}}$  threshold estimates (i.e., UT; at 95<sup>th</sup> percentile) for the target compounds, based on their  
582 CDFs (Fig. 4). Although the robustness of these values largely depends on the quality and the scope of the underlying metadata,  
583 our bootstrapping approach addresses these potential limitations. Nevertheless, we acknowledge that these estimates remain  
584 data-constrained and therefore can improve with the inclusion of more comprehensive, high-resolution datasets across diverse

585 environmental conditions. Measured concentrations beyond these thresholds must be considered exceptionally high and  
 586 warrant closer investigation to determine whether they reflect specific compounds or environmental conditions, such as  
 587 biogeochemical, oceanographic and weather-related factors. High concentrations of CHO ( $> 50 \mu\text{mol L}^{-1}$ ) were reported by  
 588 Milinković et al. (2022), which affect the outcome of the UT estimate. In contrast, typical CHO concentrations in the SML  
 589 reported in other studies remain well below the  $50 \mu\text{mol L}^{-1}$ . Since these differences cannot be resolved here, our UT estimate  
 590 for CHO should therefore be treated with caution.

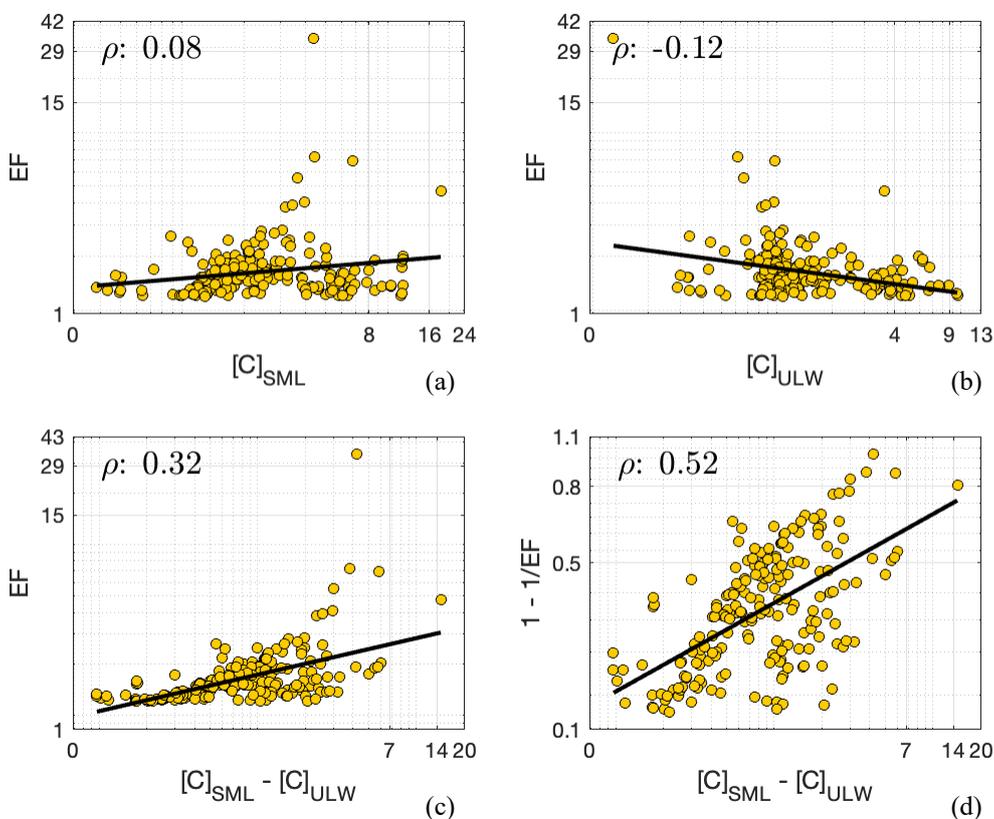
591 **Table 1: A summary of estimated UT values (upper threshold; concentration at 95<sup>th</sup> percentile) for  $[C]_{\text{SML}}$  distributions of the target**  
 592 **compounds.** This metric represents the maximum accumulation capacity of a certain compound in the SML. The values are rounded to the  
 593 nearest whole number.

Compound	TOC mg L <sup>-1</sup>	POC mg L <sup>-1</sup>	DOC mg L <sup>-1</sup>	TON mg L <sup>-1</sup>	PON mg L <sup>-1</sup>	DON mg L <sup>-1</sup>	AA $\mu\text{mol L}^{-1}$	FA $\mu\text{g L}^{-1}$	TEP $\mu\text{g Xeq L}^{-1}$	CHO $\mu\text{mol L}^{-1}$	Lipids $\mu\text{mol L}^{-1}$	Proteins $\mu\text{mol L}^{-1}$
UT value	14.0	11.0	10.0	2.0	1.0	0.5	6.0	46.0	6166.0	56.0	8.0	30.0

594 Considering absolute changes in the SML concentrations (rather than relative changes) – calculated as the magnitude difference  
 595 between corresponding SML and ULW concentrations (i.e.,  $[C]_{\text{SML}} - [C]_{\text{ULW}}$ ) – provides complementary insights into the  
 596 SML’s enrichment dynamics. When this metric is compared against the EF values for DOC data where  $\text{EF} > \bar{x}$  ( $= 1.2$ ), resulting  
 597 Spearman’s correlation coefficients ( $\rho$ ) reveal a stronger relationship (Figure 10(c);  $\rho = 0.32$ ) relative to EF vs.  $[C]_{\text{SML}}$  (Figure  
 598 10(a);  $\rho = 0.08$ ) and EF vs.  $[C]_{\text{ULW}}$  (Figure 10(b);  $\rho = -0.12$ ) correlations. This implies that although the EFs may have a limited  
 599 capacity to represent the absolute concentrations of either SML or ULW, they are more responsive to the absolute concentration  
 600 ‘changes’ in the two compartments. This analysis reveals that although ‘enrichment factor’ obscures accurately interpreting  
 601 the trophic status or the actual enrichment in the SML, it may still hold value as a proxy that reflects the degree of partitioning  
 602 between the surface microlayer and underlying waters.

603 Furthermore, normalization of  $[C]_{\text{SML}} - [C]_{\text{ULW}}$  metric to the corresponding  $[C]_{\text{SML}}$  values (i.e.,  $([C]_{\text{SML}} - [C]_{\text{ULW}})/[C]_{\text{SML}}$ )  
 604 ultimately yields an EF-based metric:  $1 - \frac{1}{\text{EF}}$ . This expresses how much of the SML concentration is above the ULW baseline,  
 605 effectively providing a measure of fractional enrichment that overlooks background variability in the ULW. Unlike  
 606 conventional EF values,  $1 - \frac{1}{\text{EF}}$  only ranges between 0 and 1. It rescales compound-specific variability in EF and is therefore  
 607 better suited for comparison across all the different observational types; normalization of EF onto a common scale allows  
 608 direct evaluations without bias from different units, magnitudes or concentration ranges. This metric better captures true trends,  
 609 rather than artifacts/effects of scale, while enhancing visualization and communication of results. In addition, when compared  
 610 against the absolute changes, this metric exhibits stronger correlation ( $\rho = 0.52$ ; Fig. 10(d)), likely due to increased robustness  
 611 to concentration variability obtained through its scale-dependent nature. As a result, when incorporated into modelling efforts,

612 the normalized EF metric can offer distinct advantages such as integration of heterogenous datasets, consistent  
613 parameterization, easier comparisons of model predictions and robust sensitivity analyses. Together, these benefits contribute  
614 to more reliable and generalizable models of SML processes.



615 **Figure 10: Correlations between EF and (a)  $[C]_{\text{SML}}$ , (b)  $[C]_{\text{ULW}}$ , (c)  $[C]_{\text{SML}} - [C]_{\text{ULW}}$  and, (d) correlation between  $1 - 1/\text{EF}$  and  $[C]_{\text{SML}} - [C]_{\text{ULW}}$ .**  
616 **These plots were generated for DOC data, in order to investigate the observed lack of correlation among  $\text{EF} - [C]_{\text{SML}} - [C]_{\text{ULW}}$ ,**  
617 **as shown by Fig. 7(c).**

## 618 5. Conclusion

619 This study presents the first known meta-analysis of the SML, integrating a broad dataset of  $[C]_{\text{SML}}$  and  $[C]_{\text{ULW}}$  measurements  
620 to resolve methodological inconsistencies and establish a consensus-based understanding of SML enrichment dynamics. By  
621 meeting the statistical requirements for combining EF data and applying KDE as a robust analytical framework, we provide  
622 reliable distributional estimates and redefine typical EF ranges for 12 organic compounds, offering a comprehensive reference  
623 for assessing whether new observations fall within expected conditions or reflect unusual enrichment. Our results indicate that  
624 nitrogen-rich compounds and particulate OM exhibit stronger enrichment than carbon-rich and dissolved compounds.

625 Nevertheless, the differing enrichment behavior of individual surfactants highlights that their surface-active properties, rather  
626 than elemental composition alone, govern overall SML enrichment. Amongst these, the fatty acids clearly show the greatest  
627 potential for high enrichment in the SML. This emphasizes the need to consider compound-specific chemistry as well as  
628 environmental and methodological variability when interpreting SML processes, assessing their role in global gas flux  
629 estimates and, developing models. Our assessment also inquired into the suitability of EF values as indicators of true SML  
630 enrichment and suggests that, while EFs capture relative partitioning between the SML and ULW, they fall short in resolving  
631 trophic variability. This study proposes complementary metrics (i.e., absolute concentration differences, SML concentration  
632 capacities and fractional enrichment) that isolate true SML enrichment and support improved SML modelling. Finally, we  
633 demonstrate that logarithmic transformations and robust central tendency metrics substantially improve statistical reliability  
634 and data comparability over traditional linear-scale approaches, providing essential methodological guidance for future SML  
635 research and its application to global air-sea exchange studies.

### 636 **Appendix A: KDE method – additional information**

637 Although the most basic non-parametric method to derive a probability distribution is histograms, they present two key  
638 limitations for comparative studies: (1) unequal sample sizes across comparative groups restrict the use of uniform binning  
639 and, (2) imposing uniform bin sizes potentially mask important distributional characteristics. In contrast, KDE circumvents  
640 these issues by accounting a datapoint's exact value rather than assigning it to a particular bin of a certain width. This describes  
641 the true underlying distribution of the data and allows more consistent and detailed comparisons of distributions. In this  
642 analysis, we use Gaussian kernels – smooth, bell-shaped functions based on normal distribution – that weight observations  
643 based on their distance. Chen (2017) and the references therein provide a comprehensive review of the KDE and its recent  
644 advances.

645 In Gaussian kernels, bandwidth is analogous to standard deviation. In this study, the bandwidths for the linear KDEs were  
646 computed based on an approach that includes a bias-variance trade-off. Briefly, the bias-variance trade-off represents kernels  
647 that have a bandwidth that avoids too much variance in the estimates (i.e., bandwidths are not too small) while it does not  
648 introduce too much bias for ranges that actually exhibit no data points (i.e., bandwidths are not too large). Calculations of  
649 optimal bandwidth applied herein and an example of a bias-variance trade-off are described in Schartau *et al.*, (2010).  
650 Nevertheless, in log-space, unlike in linear-space, data are more evenly distributed and hence fixed bandwidths avoid over-  
651 smoothing of low values and under-smoothing of high values.

652 The selection of an optimal bandwidth for KDEs is influenced by sample size; smaller sample sizes lead to sparse and noisy  
653 distributions which require more smoothing and therefore larger bandwidths. Excessively large bandwidths can result in  
654 underfitting. In contrast, larger sample sizes may allow excessively smaller bandwidths that can lead to overfitting. Bootstrap

655 resampling addresses these potential uncertainties in our analysis and, ensures the robustness and precision of the estimated  
 656 density distributions. Deviations between the bootstrapped KDEs and their ensemble mean were found to approximate a  
 657 normal distribution (consistent with Central Limit Theorem). Therefore, these ensemble means can be regarded as reliable  
 658 representations of the underlying data, supporting valid comparisons of probability distributions across different groups or  
 659 clusters.

## 660 **Appendix B: Mathematical expressions of distributional characteristics**

661 If a dataset contains values of ‘ $x_i$ ’ with a sample size of ‘ $n$ ’, mode ( $x_m$ ) is the most frequently occurring value in the dataset  
 662 and therefore, the point where a PDF reaches its highest density. A distribution appears to be the most concentrated at  $x_m$ .  
 663 Median ( $\tilde{x}$ ) returns the value at the 50<sup>th</sup> percentile of an ascending dataset. It divides the area under a PDF into two equal  
 664 halves. The arithmetic mean ( $\overline{x}_a$ ), is the average of a distribution, given by the following equation:

$$665 \quad \sum_{i=1}^n \frac{x_i}{n} \quad (B1)$$

666  $\overline{x}_a$  gives the point where weighted sum of a PDF is balanced. However, in the case of datasets that range over several orders  
 667 of magnitudes, the geometric mean (hereafter referred to as ‘ $\overline{x}_g$ ’) is the more preferred central tendency estimate, as it accounts  
 668 for the relative proportions of values (as opposed to their absolute magnitudes as is the case in  $\overline{x}_a$ ) and hence, is less sensitive  
 669 to outliers.  $\overline{x}_g$  is calculated by the following equation:

$$670 \quad (\prod_{i=1}^n x_i)^{\frac{1}{n}} \quad (B2)$$

671  $\overline{x}_g$  of a linear distribution is mathematically equal to the exponentiated  $\overline{x}_a$  of the log-transformed version of the same  
 672 distribution.

673 In addition, the following equation, which accounts for the squared differences across all the datapoints of the corresponding  
 674 CDFs, estimates the discrete form of the integrated quadratic distance (IQD, explained in section 2.2.2), with  $\Delta x_i = x_i - x_{i-1}$ :

$$675 \quad IQD = \sum_{i=1}^n \left( \left( CDF_{[C]_{SML}(x_i)} - CDF_{[C]_{ULW}(x_i)} \right)^2 \times \Delta x_i \right) \quad (B3)$$

676 A higher IQD value implies that the divergence is greater and therefore the corresponding CDFs are more different.

## 677 **Code availability**

678 Computational codes used in this study are available at OceanRep GEOMAR [ <https://oceanrep.geomar.de/id/eprint/63615/> ].  
 679 The repository includes the implementation of the KDE method, representative example scripts demonstrating its application  
 680 for generating probability density functions and cumulative density functions, and a script to reproduce the correlation plots  
 681 presented in the manuscript.

682 **Data availability**

683 All data used in this study were extracted from previously published peer-reviewed sources and are publicly available through  
684 the PANGAEA data repository [ <https://doi.pangaea.de/10.1594/PANGAEA.990017> ]. Full citations for all the datasets are  
685 provided in supplementary information. No new data were generated for this study.

686 **Author contribution**

687 A. S. – Data Curation, Conceptualization, Methodology, Formal Analysis, Visualization, Writing – Original Draft, Writing –  
688 Review & Editing

689 S. N. – Data Curation, Writing – Review & Editing

690 T. B. – Data Provision, Writing – Review & Editing

691 A. E. – Funding Acquisition, Data Provision, Writing – Review & Editing

692 H. H. – Data Provision, Writing – Review & Editing

693 M. P. – Data Provision, Writing – Review & Editing

694 K. W. – Methodology, Writing – Review & Editing

695 O. W. – Funding Acquisition, Data Provision, Writing – Review & Editing

696 M. S. – Conceptualization, Methodology, Formal Analysis, Visualization, Funding Acquisition, Supervision, Writing –  
697 Review & Editing

698 **Competing interests**

699 Authors A. S., T. B., A. E. and M. S. are affiliated with the same institution as H. B., who serves as an overseeing editor for  
700 the special issue “*Biogeochemical processes and air-sea exchange in the sea-surface microlayer*”. Authors A. S., T. B., A.  
701 E., H. H., M. P., O. W. and M. S. are collaborators with H. B. on an ongoing research project. These potential competing  
702 interests have been fully disclosed to the journal. The authors declare no other competing interests relevant to the submitted  
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## 714 **References**

- 715 Alldredge, A.L., Passow, U., Logan, B.E., 1993. The abundance and significance of a class of large, transparent organic  
716 particles in the ocean. *Deep Sea Res. Part Oceanogr. Res. Pap.* 40, 1131–1140. [https://doi.org/10.1016/0967-](https://doi.org/10.1016/0967-0637(93)90129-Q)  
717 [0637\(93\)90129-Q](https://doi.org/10.1016/0967-0637(93)90129-Q)
- 718 Asher, W., 1997. The sea-surface microlayer and its effect on global air-sea gas transfer. *Sea Surf. Glob. Change* 251–286.
- 719 Asmussen-Schäfer, F., Ribas-Ribas, M., Wurl, O., Friedrichs, G., 2026. Linking surface coverage with surfactant activity to  
720 refine the role of surfactants for air-sea gas exchange. <https://doi.org/10.5194/egusphere-2025-5276>
- 721 Astrahan, P., Herut, B., Paytan, A., Rahav, E., 2016. The Impact of Dry Atmospheric Deposition on the Sea-Surface Microlayer  
722 in the SE Mediterranean Sea: An Experimental Approach. *Front. Mar. Sci.* 3.  
723 <https://doi.org/10.3389/fmars.2016.00222>
- 724 Baastrup-Spohr, L., Staehr, P.A., 2009. Surface microlayers on temperate lowland lakes. *Hydrobiologia* 625, 43–59.  
725 <https://doi.org/10.1007/s10750-008-9695-3>
- 726 Baier, R., DW, G., Perlmutter, S., King, R., 1974. Dominant chemical composition of sea-surface films, natural slicks, and  
727 foams.
- 728 Barthelmeß, T., Engel, A., 2022. How biogenic polymers control surfactant dynamics in the surface microlayer: insights from  
729 a coastal Baltic Sea study. *Biogeosciences* 19, 4965–4992. <https://doi.org/10.5194/bg-19-4965-2022>
- 730 Barthelmeß, T., Schütte, F., Engel, A., 2021. Variability of the Sea Surface Microlayer Across a Filament’s Edge and Potential  
731 Influences on Gas Exchange. *Front. Mar. Sci.* 8, 718384. <https://doi.org/10.3389/fmars.2021.718384>
- 732 Brinis, A., Méjanelle, L., Momzikoff, A., Gondry, G., Fillaux, J., Point, V., Saliot, A., 2004. Phospholipid ester-linked fatty  
733 acids composition of size-fractionated particles at the top ocean surface. *Org. Geochem.* 35, 1275–1287.  
734 <https://doi.org/10.1016/j.orggeochem.2004.04.009>
- 735 Carlson, D.J., 1983. Dissolved organic materials in surface microlayers: Temporal and spatial variability and relation to sea  
736 state. *Limnol. Oceanogr.* 28, 415–431. <https://doi.org/10.4319/lo.1983.28.3.0415>
- 737 Carlucci, A.F., Craven, D.B., Henrichs, S.M., 1985. Surface-film microheterotrophs: amino acid metabolism and solar  
738 radiation effects on their activities.
- 739 Cen-Lin, H., Tzung-May, F., 2013. Air-Sea Exchange of Volatile Organic Compounds: A New Model with Microlayer Effects.  
740 *Atmospheric Ocean. Sci. Lett.* 6, 97–102. <https://doi.org/10.1080/16742834.2013.11447063>
- 741 Chen, Y., Yang, G.-P., Xia, Q.-Y., Wu, G.-W., 2016. Enrichment and characterization of dissolved organic matter in the  
742 surface microlayer and subsurface water of the South Yellow Sea. *Mar. Chem.* 182, 1–13.  
743 <https://doi.org/10.1016/j.marchem.2016.04.001>
- 744 Chen, Y.-C., 2017. A tutorial on kernel density estimation and recent advances. *Biostat. Epidemiol.* 1, 161–187.
- 745 Čosović, B., Vojvodić, V., 1998. Voltammetric Analysis of Surface Active Substances in Natural Seawater. *Electroanalysis*  
746 10, 429–434. [https://doi.org/10.1002/\(SICI\)1521-4109\(199805\)10:6%253C429::AID-ELAN429%253E3.0.CO;2-7](https://doi.org/10.1002/(SICI)1521-4109(199805)10:6%253C429::AID-ELAN429%253E3.0.CO;2-7)
- 747 Čosović, B., Vojvodić, V., 1989. Adsorption behaviour of the hydrophobic fraction of organic matter in natural waters. *Mar.*  
748 *Chem.* 28, 183–198. [https://doi.org/10.1016/0304-4203\(89\)90194-1](https://doi.org/10.1016/0304-4203(89)90194-1)
- 749 Crocetti, E., 2016. Systematic Reviews With Meta-Analysis: Why, When, and How? *Emerg. Adulthood* 4, 3–18.  
750 <https://doi.org/10.1177/2167696815617076>

- 751 Cunliffe, M., Engel, A., Frka, S., Gašparović, B., Guitart, C., Murrell, J.C., Salter, M., Stolle, C., Upstill-Goddard, R., Wurl,  
752 O., 2013. Sea surface microlayers: A unified physicochemical and biological perspective of the air–ocean interface.  
753 *Prog. Oceanogr.* 109, 104–116. <https://doi.org/10.1016/j.pocean.2012.08.004>
- 754 Cunliffe, M., Murrell, J.C., 2009. The sea-surface microlayer is a gelatinous biofilm. *ISME J.* 3, 1001–1003.  
755 <https://doi.org/10.1038/ismej.2009.69>
- 756 Cunliffe, M., Salter, M., Mann, P.J., Whiteley, A.S., Upstill-Goddard, R.C., Murrell, J.C., 2009. Dissolved organic carbon and  
757 bacterial populations in the gelatinous surface microlayer of a Norwegian fjord mesocosm. *FEMS Microbiol. Lett.*  
758 299, 248–254. <https://doi.org/10.1111/j.1574-6968.2009.01751.x>
- 759 Dietz, A.S., Albright, L.J., Tuominen, T., 1976. Heterotrophic activities of bacterioneuston and bacterioplankton. *Can. J.*  
760 *Microbiol.* 22, 1699–1709. <https://doi.org/10.1139/m76-251>
- 761 Dukhin, S.S., Kovalchuk, V.I., Gochev, G.G., Lotfi, M., Krzan, M., Malysa, K., Miller, R., 2015. Dynamics of Rear Stagnant  
762 Cap formation at the surface of spherical bubbles rising in surfactant solutions at large Reynolds numbers under  
763 conditions of small Marangoni number and slow sorption kinetics. *Adv. Colloid Interface Sci.* 222, 260–274.  
764 <https://doi.org/10.1016/j.cis.2014.10.002>
- 765 Engel, A., Bange, H.W., Cunliffe, M., Burrows, S.M., Friedrichs, G., Galgani, L., Herrmann, H., Hertkorn, N., Johnson, M.,  
766 Liss, P.S., Quinn, P.K., Schartau, M., Soloviev, A., Stolle, C., Upstill-Goddard, R.C., Van Pinxteren, M., Zäncker,  
767 B., 2017. The Ocean’s Vital Skin: Toward an Integrated Understanding of the Sea Surface Microlayer. *Front. Mar.*  
768 *Sci.* 4, 165. <https://doi.org/10.3389/fmars.2017.00165>
- 769 Engel, A., Galgani, L., 2016. The organic sea-surface microlayer in the upwelling region off the coast of Peru and potential  
770 implications for air–sea exchange processes. *Biogeosciences* 13, 989–1007. <https://doi.org/10.5194/bg-13-989-2016>
- 771 Engel, A., Thoms, S., Riebesell, U., Rochelle-Newall, E., Zondervan, I., 2004. Polysaccharide aggregation as a potential sink  
772 of marine dissolved organic carbon. *Nature* 428, 929–932. <https://doi.org/10.1038/nature02453>
- 773 Feenstra, S., 2006. Use of logarithmic-scale correlation plots to represent contaminant ratios for evaluation of subsurface  
774 environmental data. *Environ. Forensics* 7, 175–185.
- 775 Frew, N.M., 1997. The role of organic films in air–sea gas exchange, in: Liss, P.S., Duce, R.A. (Eds.), *The Sea Surface and*  
776 *Global Change*. Cambridge University Press, Cambridge, pp. 121–172.  
777 <https://doi.org/10.1017/CBO9780511525025.006>
- 778 Frew, N.M., Bock, E.J., Schimpf, U., Hara, T., Haußecker, H., Edson, J.B., McGillis, W.R., Nelson, R.K., McKenna, S.P., Uz,  
779 B.M., Jähne, B., 2004. Air–sea gas transfer: Its dependence on wind stress, small-scale roughness, and surface films.  
780 *J. Geophys. Res.* Oceans 109, 2003JC002131. <https://doi.org/10.1029/2003JC002131>
- 781 Frka, S., Pogorzelski, S., Kozarac, Z., Čosović, B., 2012. Physicochemical Signatures of Natural Sea Films from Middle  
782 Adriatic Stations. *J. Phys. Chem. A* 116, 6552–6559. <https://doi.org/10.1021/jp212430a>
- 783 Gao, Q., Leck, C., Rauschenberg, C., Matrai, P.A., 2012. On the chemical dynamics of extracellular polysaccharides in the  
784 high Arctic surface microlayer. *Ocean Sci.* 8, 401–418. <https://doi.org/10.5194/os-8-401-2012>
- 785 Garabetian, F., Romano, J.-C., Paul, R., Sigoillot, J.-C., 1993. Organic matter composition and pollutant enrichment of sea  
786 surface microlayer inside and outside slicks. *Mar. Environ. Res.* 35, 323–339. [https://doi.org/10.1016/0141-1136\(93\)90100-E](https://doi.org/10.1016/0141-1136(93)90100-E)
- 787
- 788 Gašparović, B., Čosović, B., 2003. Surface-active properties of organic matter in the North Adriatic Sea. *Estuar. Coast. Shelf*  
789 *Sci.* 58, 555–566. [https://doi.org/10.1016/S0272-7714\(03\)00133-1](https://doi.org/10.1016/S0272-7714(03)00133-1)
- 790 Gašparović, B., Čosović, B., 2001. Distribution of surface-active substances in the northern Adriatic Sea. *Mar. Chem.* 75, 301–  
791 313. [https://doi.org/10.1016/S0304-4203\(01\)00044-5](https://doi.org/10.1016/S0304-4203(01)00044-5)
- 792 Gašparović, B., Plavšić, M., Čosović, B., Saliot, A., 2007. Organic matter characterization in the sea surface microlayers in  
793 the subarctic Norwegian fjords region. *Mar. Chem.* 105, 1–14. <https://doi.org/10.1016/j.marchem.2006.12.010>
- 794 Goldman, J.C., Dennett, M.R., Frew, N.M., 1988. Surfactant effects on air–sea gas exchange under turbulent conditions. *Deep*  
795 *Sea Res. Part Oceanogr. Res. Pap.* 35, 1953–1970. [https://doi.org/10.1016/0198-0149\(88\)90119-7](https://doi.org/10.1016/0198-0149(88)90119-7)
- 796 Härdle, W., Müller, M., Sperlich, S., Werwatz, A., 2004. *Nonparametric and semiparametric models*. Springer Science &  
797 *Business Media*.
- 798 Harvey, H.R., Tuttle, J.H., Bell, J.T., 1995. Kinetics of phytoplankton decay during simulated sedimentation: Changes in  
799 biochemical composition and microbial activity under oxic and anoxic conditions. *Geochim. Cosmochim. Acta* 59,  
800 3367–3377. [https://doi.org/10.1016/0016-7037\(95\)00217-N](https://doi.org/10.1016/0016-7037(95)00217-N)

- 801 Hedges, J.I., Cowie, G.L., Richey, J.E., Quay, P.D., Benner, R., Strom, M., Forsberg, B.R., 1994. Origins and processing of  
802 organic matter in the Amazon River as indicated by carbohydrates and amino acids. *Limnol. Oceanogr.* 39, 743–761.  
803 <https://doi.org/10.4319/lo.1994.39.4.0743>
- 804 Henrichs, S.M., Williams, P.M., 1985. Dissolved and particulate amino acids and carbohydrates in the sea surface microlayer.  
805 *Mar. Chem.* 17, 141–163. [https://doi.org/10.1016/0304-4203\(85\)90070-2](https://doi.org/10.1016/0304-4203(85)90070-2)
- 806 Hillbricht-Ilkowska, A., Kostrzewska-Szlakowska, I., 2004. Surface microlayer in lakes of different trophic status: nutrients  
807 concentration and accumulation. *Pol. J. Ecol.* 52, 461–478.
- 808 Hunter, K.A., 1980. Processes affecting particulate trace metals in the sea surface microlayer. *Mar. Chem.* 9, 49–70.  
809 [https://doi.org/10.1016/0304-4203\(80\)90006-7](https://doi.org/10.1016/0304-4203(80)90006-7)
- 810 Hunter, K.A., Liss, P.S., 1981. Chapter 9 Organic Sea Surface Films, in: Duursma, E.K., Dawson, R. (Eds.), Elsevier  
811 Oceanography Series. Elsevier, pp. 259–298. [https://doi.org/10.1016/S0422-9894\(08\)70331-3](https://doi.org/10.1016/S0422-9894(08)70331-3)
- 812 Hunter, K.A., Liss, P.S., 1977. The input of organic material to the oceans: air–sea interactions and the organic chemical  
813 composition of the sea surface. *Mar. Chem.* 5, 361–379. [https://doi.org/10.1016/0304-4203\(77\)90029-9](https://doi.org/10.1016/0304-4203(77)90029-9)
- 814 Jørgensen, N., Kroer, N., Coffin, R., Yang, X.-H., Lee, C., 1993. Dissolved free amino acids, combined amino acids, and DNA  
815 as sources of carbon and nitrogen to marine bacteria. *Mar. Ecol. Prog. Ser.* 98, 135–148.  
816 <https://doi.org/10.3354/meps098135>
- 817 Joux, F., Agogué, H., Obernosterer, I., Dupuy, C., Reinthaler, T., Herndl, G., Lebaron, P., 2006. Microbial community structure  
818 in the sea surface microlayer at two contrasting coastal sites in the northwestern Mediterranean Sea. *Aquat. Microb.  
819 Ecol.* 42, 91–104. <https://doi.org/10.3354/ame042091>
- 820 Keene, O.N., 1995. The log transformation is special. *Stat. Med.* 14, 811–819. <https://doi.org/10.1002/sim.4780140810>
- 821 Kjelleberg, S., Norkrans, B., Löfgren, H., Larsson, K., 1976. Surface balance study of the interaction between microorganisms  
822 and lipid monolayer at the air/water interface. *Appl. Environ. Microbiol.* 31, 609–611.  
823 <https://doi.org/10.1128/aem.31.4.609-611.1976>
- 824 Knipping, E.M., Lakin, M.J., Foster, K.L., Jungwirth, P., Tobias, D.J., Gerber, R.B., Dabdub, D., Finlayson-Pitts, B.J., 2000.  
825 Experiments and Simulations of Ion-Enhanced Interfacial Chemistry on Aqueous NaCl Aerosols. *Science* 288, 301–  
826 306. <https://doi.org/10.1126/science.288.5464.301>
- 827 Knulst, J.C., Backlund, P., Hessen, D.O., Riise, G., Södergren, A., 1997. Response of surface microlayers to artificial acid  
828 precipitation in a meso-humic lake in Norway. *Water Res.* 31, 2177–2186. [https://doi.org/10.1016/S0043-1354\(97\)00061-4](https://doi.org/10.1016/S0043-1354(97)00061-4)
- 830 Kock, A., Schafstall, J., Dengler, M., Brandt, P., Bange, H.W., 2012. Sea-to-air and diapycnal nitrous oxide fluxes in the  
831 eastern tropical North Atlantic Ocean. *Biogeosciences* 9, 957–964. <https://doi.org/10.5194/bg-9-957-2012>
- 832 Kujawinski, E.B., Farrington, J.W., Moffett, J.W., 2002. Evidence for grazing-mediated production of dissolved surface-active  
833 material by marine protists. *Mar. Chem.* 77, 133–142. [https://doi.org/10.1016/S0304-4203\(01\)00082-2](https://doi.org/10.1016/S0304-4203(01)00082-2)
- 834 Kurata, N., Vella, K., Hamilton, B., Shivji, M., Soloviev, A., Matt, S., Tartar, A., Perrie, W., 2016. Surfactant-associated  
835 bacteria in the near-surface layer of the ocean. *Sci. Rep.* 6, 19123. <https://doi.org/10.1038/srep19123>
- 836 Kuznetsova, M., Lee, C., 2002. Dissolved free and combined amino acids in nearshore seawater, sea surface microlayers and  
837 foams: Influence of extracellular hydrolysis. *Aquat Sci* 64, 252–268.
- 838 Kuznetsova, M., Lee, C., Aller, J., Frew, N., 2004. Enrichment of amino acids in the sea surface microlayer at coastal and open  
839 ocean sites in the North Atlantic Ocean. *Limnol. Oceanogr.* 49, 1605–1619.  
840 <https://doi.org/10.4319/lo.2004.49.5.1605>
- 841 Laß, K., Bange, H.W., Friedrichs, G., 2013. Seasonal signatures in SFG vibrational spectra of the sea surface nanolayer at  
842 Boknis Eck Time Series Station (SW Baltic Sea). *Biogeosciences* 10, 5325–5334. <https://doi.org/10.5194/bg-10-5325-2013>
- 843
- 844 Laß, K., Friedrichs, G., 2011. Revealing structural properties of the marine nanolayer from vibrational sum frequency  
845 generation spectra. *J. Geophys. Res.* 116, C08042. <https://doi.org/10.1029/2010JC006609>
- 846 Li, L., Wu, P., Zhang, P., Huang, S., Zhang, Y., 2024. An improved model for air–sea exchange of elemental mercury in  
847 MITgcm-ECCOv4-Hg: the role of surfactants and waves. *Geosci. Model Dev.* 17, 8683–8695.  
848 <https://doi.org/10.5194/gmd-17-8683-2024>
- 849 Liss, P.S., Duce, R.A., 1997. The sea surface and global change. Cambridge University Press.

- 850 Liu, K., Dickhut, R.M., 1998. Effects of wind speed and particulate matter source on surface microlayer characteristics and  
851 enrichment of organic matter in southern Chesapeake Bay. *J. Geophys. Res. Atmospheres* 103, 10571–10577.  
852 <https://doi.org/10.1029/97JD03736>
- 853 López-Puertas, A., Wurl, O., Frka, S., Ribas-Ribas, M., 2025. Diel Variability Affects the Inorganic Marine Carbon System  
854 in the Sea-Surface Microlayer of a Mediterranean coastal area (Šibenik, Croatia). <https://doi.org/10.5194/egusphere-2025-2090>
- 855
- 856 MacIntyre, F., 1974. Chemical fractionation and sea-surface microlayer processes. *The sea* 5, 245–299.
- 857 Maki, J.S., Hermansson, M., 2020. The dynamics of surface microlayers in aquatic environments. *Biol. Part. Aquat. Syst.*  
858 Second Ed. 161–182.
- 859 Mari, X., Burd, A., 1998. Seasonal size spectra of transparent exopolymeric particles (TEP) in a coastal sea and comparison  
860 with those predicted using coagulation theory. *Mar. Ecol. Prog. Ser.* 163, 63–76. <https://doi.org/10.3354/meps163063>
- 861 Mari, X., Passow, U., Migon, C., Burd, A.B., Legendre, L., 2017. Transparent exopolymer particles: Effects on carbon cycling  
862 in the ocean. *Prog. Oceanogr.* 151, 13–37. <https://doi.org/10.1016/j.pocean.2016.11.002>
- 863 Marty, J.C., Saliot, A., 1976. Hydrocarbons (normal alkanes) in the surface microlayer of seawater. *Deep-Sea Res.* 23, 863–  
864 873.
- 865 McKenna, S.P., McGillis, W.R., 2004. The role of free-surface turbulence and surfactants in air–water gas transfer. *Int. J. Heat*  
866 *Mass Transf.* 47, 539–553. <https://doi.org/10.1016/j.ijheatmasstransfer.2003.06.001>
- 867 Menge, D.N.L., MacPherson, A.C., Bytnerowicz, T.A., Quebbeman, A.W., Schwartz, N.B., Taylor, B.N., Wolf, A.A., 2018.  
868 Logarithmic scales in ecological data presentation may cause misinterpretation. *Nat. Ecol. Evol.* 2, 1393–1402.  
869 <https://doi.org/10.1038/s41559-018-0610-7>
- 870 Mengist, W., Soromessa, T., Legese, G., 2020. Method for conducting systematic literature review and meta-analysis for  
871 environmental science research. *MethodsX* 7, 100777. <https://doi.org/10.1016/j.mex.2019.100777>
- 872 Milinković, A., Penezić, A., Kušan, A.C., Gluščić, V., Žužul, S., Skejić, S., Šantić, D., Godec, R., Pehnc, G., Omanović, D.,  
873 Engel, A., Frka, S., 2022. Variabilities of biochemical properties of the sea surface microlayer: Insights to the  
874 atmospheric deposition impacts. *Sci. Total Environ.* 838, 156440. <https://doi.org/10.1016/j.scitotenv.2022.156440>
- 875 Mopper, K., Zhou, J., Sri Ramana, K., Passow, U., Dam, H.G., Drapeau, D.T., 1995. The role of surface-active carbohydrates  
876 in the flocculation of a diatom bloom in a mesocosm. *Deep Sea Res. Part II Top. Stud. Oceanogr.* 42, 47–73.  
877 [https://doi.org/10.1016/0967-0645\(95\)00004-A](https://doi.org/10.1016/0967-0645(95)00004-A)
- 878 Münster, U., Heikkinen, E., Knulst, J., 1998. Nutrient composition, microbial biomass and activity at the air-water interface  
879 of small boreal forest lakes, in: Tamminen, T., Kuosa, H. (Eds.), *Eutrophication in Planktonic Ecosystems: Food Web*  
880 *Dynamics and Elemental Cycling*. Springer Netherlands, Dordrecht, pp. 261–270. [https://doi.org/10.1007/978-94-017-1493-8\\_21](https://doi.org/10.1007/978-94-017-1493-8_21)
- 881
- 882 Mustaffa, N.I.H., Badewien, T.H., Ribas-Ribas, M., Wurl, O., 2018. High-resolution observations on enrichment processes in  
883 the sea-surface microlayer. *Sci. Rep.* 8, 13122. <https://doi.org/10.1038/s41598-018-31465-8>
- 884 Mustaffa, N.I.H., Ribas-Ribas, M., Banko-Kubis, H.M., Wurl, O., 2020. Global reduction of *in situ* CO<sub>2</sub> transfer velocity by  
885 natural surfactants in the sea-surface microlayer. *Proc. R. Soc. Math. Phys. Eng. Sci.* 476, 20190763.  
886 <https://doi.org/10.1098/rspa.2019.0763>
- 887 Myklestad, S.M., 1995. Release of extracellular products by phytoplankton with special emphasis on polysaccharides. *Sci.*  
888 *Total Environ.* 165, 155–164. [https://doi.org/10.1016/0048-9697\(95\)04549-G](https://doi.org/10.1016/0048-9697(95)04549-G)
- 889 Obernosterer, I., Catala, P., Lami, R., Caparros, J., Ras, J., Bricaud, A., Dupuy, C., van Wambeke, F., Lebaron, P., 2008.  
890 Biochemical characteristics and bacterial community structure of the sea surface microlayer in the South Pacific  
891 Ocean.
- 892 Obernosterer, I., Catala, P., Reinthaler, T., Herndl, G., Lebaron, P., 2005. Enhanced heterotrophic activity in the surface  
893 microlayer of the Mediterranean Sea. *Aquat. Microb. Ecol.* 39, 293–302. <https://doi.org/10.3354/ame039293>
- 894 Parzen, E., 1962. On Estimation of a Probability Density Function and Mode. *Ann. Math. Stat.* 33, 1065–1076.
- 895 Passow, U., 2000. Formation of transparent exopolymer particles, TEP, from dissolved precursor material. *Mar. Ecol. Prog.*  
896 *Ser.* 192, 1–11. <https://doi.org/10.3354/meps192001>
- 897 Passow, U., Alldredge, A.L., 1999. Do transparent exopolymer particles (TEP) inhibit grazing by the euphausiid *Euphausia*  
898 *pacifica*? *J. Plankton Res.* 21, 2203–2217. <https://doi.org/10.1093/plankt/21.11.2203>

- 899 Penezić, A., Drozdowska, V., Novak, T., Gašparović, B., 2022. Distribution and characterization of organic matter within the  
900 sea surface microlayer in the Gulf of Gdańsk. *Oceanologia* 64, 631–650.  
901 <https://doi.org/10.1016/j.oceano.2022.05.003>
- 902 Penna, A., 1999. Influence of nutrient ratios on the in vitro extracellular polysaccharide production by marine diatoms from  
903 the Adriatic Sea. *J. Plankton Res.* 21, 1681–1690. <https://doi.org/10.1093/plankt/21.9.1681>
- 904 Pereira, R., Ashton, I., Sabbaghzadeh, B., Shutler, J.D., Upstill-Goddard, R.C., 2018. Reduced air–sea CO<sub>2</sub> exchange in the  
905 Atlantic Ocean due to biological surfactants. *Nat. Geosci.* 11, 492–496. <https://doi.org/10.1038/s41561-018-0136-2>
- 906 Pereira, R., Schneider-Zapp, K., Upstill-Goddard, R.C., 2016. Surfactant control of gas transfer velocity along an offshore  
907 coastal transect: results from a laboratory gas exchange tank. *Biogeosciences* 13, 3981–3989.  
908 <https://doi.org/10.5194/bg-13-3981-2016>
- 909 Petersen, M.K., Iyengar, S.S., Day, T.J.F., Voth, G.A., 2004. The Hydrated Proton at the Water Liquid/Vapor Interface. *J.*  
910 *Phys. Chem. B* 108, 14804–14806. <https://doi.org/10.1021/jp0467160>
- 911 Pogorzelski, S.J., Kogut, A.D., Mazurek, A.Z., 2006. Surface Rheology Parameters of Source-Specific Surfactant Films as  
912 Indicators of Organic Matter Dynamics. *Hydrobiologia* 554, 67–81. <https://doi.org/10.1007/s10750-005-1007-6>
- 913 Reinthaler, T., Sintes, E., Herndl, G.J., 2008. Dissolved organic matter and bacterial production and respiration in the sea-  
914 surface microlayer of the open Atlantic and the western Mediterranean Sea. *Limnol. Oceanogr.* 53, 122–136.  
915 <https://doi.org/10.4319/lo.2008.53.1.0122>
- 916 Robinson, T., Wurl, O., Bahlmann, E., Jürgens, K., Stolle, C., 2019. Rising bubbles enhance the gelatinous nature of the air–  
917 sea interface. *Limnol. Oceanogr.* 64, 2358–2372. <https://doi.org/10.1002/lno.11188>
- 918 Sabbaghzadeh, B., Upstill-Goddard, R.C., Beale, R., Pereira, R., Nightingale, P.D., 2017. The Atlantic Ocean surface  
919 microlayer from 50°N to 50°S is ubiquitously enriched in surfactants at wind speeds up to 13 m s<sup>-1</sup>. *Geophys. Res.*  
920 *Lett.* 44, 2852–2858. <https://doi.org/10.1002/2017GL072988>
- 921 Salter, M.E., Upstill-Goddard, R.C., Nightingale, P.D., Archer, S.D., Blomquist, B., Ho, D.T., Huebert, B., Schlosser, P.,  
922 Yang, M., 2011. Impact of an artificial surfactant release on air-sea gas fluxes during Deep Ocean Gas Exchange  
923 Experiment II. *J. Geophys. Res. Oceans* 116, 2011JC007023. <https://doi.org/10.1029/2011JC007023>
- 924 Schartau, M., Engel, A., Schröter, J., Thoms, S., Völker, C., Wolf-Gladrow, D., 2007. Modelling carbon overconsumption and  
925 the formation of extracellular particulate organic carbon. *Biogeosciences* 4, 433–454. <https://doi.org/10.5194/bg-4-433-2007>
- 926
- 927 Schartau, M., Landry, M.R., Armstrong, R.A., 2010. Density estimation of plankton size spectra: a reanalysis of IronEx II  
928 data. *J. Plankton Res.* 32, 1167–1184. <https://doi.org/10.1093/plankt/fbq072>
- 929 Shine, J.P., Wallace, G.T., 1996. Flux of surface-active organic complexes of copper to the air-sea interface in coastal marine  
930 waters. *J. Geophys. Res. Oceans* 101, 12017–12026. <https://doi.org/10.1029/96JC00616>
- 931 Sieburth, J.McN., 1983. Microbiological and Organic-Chemical Processes in the Surface and Mixed Layers, in: Liss, P.S.,  
932 Slinn, W.G.N. (Eds.), *Air-Sea Exchange of Gases and Particles*. Springer Netherlands, Dordrecht, pp. 121–172.  
933 [https://doi.org/10.1007/978-94-009-7169-1\\_3](https://doi.org/10.1007/978-94-009-7169-1_3)
- 934 Sieburth, J.McN., Willis, P.-J., Johnson, K.M., Burney, C.M., Lavoie, D.M., Hinga, K.R., Caron, D.A., French, F.W., Johnson,  
935 P.W., Davis, P.G., 1976. Dissolved Organic Matter and Heterotrophic Microneuston in the Surface Microlayers of  
936 the North Atlantic. *Science* 194, 1415–1418. <https://doi.org/10.1126/science.194.4272.1415>
- 937 Silverman, B.W., 1986. *Density estimation for statistics and data analysis*. Chapman & Hall.
- 938 Södergren, A., 1987. Origin and composition of surface slicks in lakes of differing trophic status. *Limnol. Oceanogr.* 32,  
939 1307–1316. <https://doi.org/10.4319/lo.1987.32.6.1307>
- 940 Springer, T.G., Pigford, R.L., 1970. Influence of Surface Turbulence and Surfactants on Gas Transport through Liquid  
941 Interfaces. *Ind. Eng. Chem. Fundam.* 9, 458–465. <https://doi.org/10.1021/i160035a025>
- 942 Sun, C.-C., Sperling, M., Engel, A., 2018. Effect of wind speed on the size distribution of gel particles in the sea surface  
943 microlayer: insights from a wind–wave channel experiment. *Biogeosciences* 15, 3577–3589.  
944 <https://doi.org/10.5194/bg-15-3577-2018>
- 945 Thornton, D.C.O., Brooks, S.D., Chen, J., 2016. Protein and Carbohydrate Exopolymer Particles in the Sea Surface Microlayer  
946 (SML). *Front. Mar. Sci.* 3. <https://doi.org/10.3389/fmars.2016.00135>
- 947 Tsai, W., Liu, K., 2003. An assessment of the effect of sea surface surfactant on global atmosphere–ocean CO<sub>2</sub> flux. *J. Geophys.*  
948 *Res. Oceans* 108, 2000JC000740. <https://doi.org/10.1029/2000JC000740>

949 Tukey, J.W., 1977. Exploratory data analysis. Springer.  
950 Upstill-Goddard, R.C., 2006. Air–sea gas exchange in the coastal zone. *Estuar. Coast. Shelf Sci.* 70, 388–404.  
951 <https://doi.org/10.1016/j.ecss.2006.05.043>  
952 Verdugo, P., Alldredge, A.L., Azam, F., Kirchman, D.L., Passow, U., Santschi, P.H., 2004. The oceanic gel phase: a bridge in  
953 the DOM–POM continuum. *Mar. Chem.* 92, 67–85. <https://doi.org/10.1016/j.marchem.2004.06.017>  
954 Vitousek, P.M., 2004. Nutrient cycling and limitation: Hawai'i as a model system. Princeton University Press.  
955 Wegman, E.J., 1972. Nonparametric probability density estimation. *J. Stat. Comput. Simul.* 1, 225–245.  
956 <https://doi.org/10.1080/00949657208810017>  
957 Williams, P.M., Carlucci, A.F., Henrichs, S.M., Van Vleet, E.S., Horrigan, S.G., Reid, F.M.H., Robertson, K.J., 1986.  
958 Chemical and microbiological studies of sea-surface films in the Southern Gulf of California and off the West Coast  
959 of Baja California. *Mar. Chem.* 19, 17–98. [https://doi.org/10.1016/0304-4203\(86\)90033-2](https://doi.org/10.1016/0304-4203(86)90033-2)  
960 Wurl, O., Holmes, M., 2008. The gelatinous nature of the sea-surface microlayer. *Mar. Chem.* 110, 89–97.  
961 <https://doi.org/10.1016/j.marchem.2008.02.009>  
962 Wurl, O., Miller, L., Röttgers, R., Vagle, S., 2009. The distribution and fate of surface-active substances in the sea-surface  
963 microlayer and water column. *Mar. Chem.* 115, 1–9. <https://doi.org/10.1016/j.marchem.2009.04.007>  
964 Wurl, O., Stolle, C., Van Thuoc, C., The Thu, P., Mari, X., 2016. Biofilm-like properties of the sea surface and predicted  
965 effects on air–sea CO<sub>2</sub> exchange. *Prog. Oceanogr.* 144, 15–24. <https://doi.org/10.1016/j.pocean.2016.03.002>  
966 Wurl, O., Wurl, E., Miller, L., Johnson, K., Vagle, S., 2011. Formation and global distribution of sea-surface microlayers.  
967 *Biogeosciences* 8, 121–135. <https://doi.org/10.5194/bg-8-121-2011>  
968 Yang, G.-P., 1999. Dimethylsulfide enrichment in the surface microlayer of the South China Sea. *Mar. Chem.* 66, 215–224.  
969 [https://doi.org/10.1016/S0304-4203\(99\)00042-0](https://doi.org/10.1016/S0304-4203(99)00042-0)  
970 Zhang, Z., Liu, C., Liu, L., 2006. Physicochemical Studies of the Sea-Surface Microlayer. *Front. Chem. China* 1.  
971 <https://doi.org/10.1007/s11458-005-0003-8>  
972 Zhou, J., Mopper, K., Passow, U., 1998. The role of surface-active carbohydrates in the formation of transparent exopolymer  
973 particles by bubble adsorption of seawater. *Limnol. Oceanogr.* 43, 1860–1871.  
974 <https://doi.org/10.4319/lo.1998.43.8.1860>  
975 Žutić, V., Čosović, B., Marčenko, E., Bihari, N., Kršinić, F., 1981. Surfactant production by marine phytoplankton. *Mar.*  
976 *Chem.* 10, 505–520. [https://doi.org/10.1016/0304-4203\(81\)90004-9](https://doi.org/10.1016/0304-4203(81)90004-9)  
977 Zuur, A.F., Ieno, E.N., Smith, G.M. (Eds.), 2007. Linear regression, in: *Analysing Ecological Data*. Springer New York, New  
978 York, NY, pp. 49–77. [https://doi.org/10.1007/978-0-387-45972-1\\_5](https://doi.org/10.1007/978-0-387-45972-1_5)  
979