

1  
2  
3  
4  
5  
6  
7  
8  
9  
10  
11  
12  
13  
14  
15  
16  
17  
18  
19  
20

**Surface Ozone Distribution & Trends Over Ireland: Insights from long-term measurement record and source attribution modelling**

Nikhil Korhale <sup>1</sup>, Tabish Ansari <sup>2</sup>, Tim Butler <sup>2</sup>, Jurgita Ovadnevaite<sup>1</sup>, Emmanuel Chevassus<sup>1</sup>, Darius Ceburnis<sup>1</sup>, Damien Martin<sup>1</sup>, Colin D.O’Dowd<sup>1</sup>, Liz Coleman<sup>1</sup>

<sup>1</sup> School of Natural Sciences, Physics, Ryan Institute’s Centre for Climate & Air Pollution Studies, University of Galway, Galway, Ireland

<sup>2</sup> Research Institute for Sustainability (RIFS) at GFZ Helmholtz Centre for Geosciences, Potsdam, 14467, Germany

**Corresponding author :**

Name - Dr. Liz Coleman  
Email Id – liz.coleman@universityofgalway.ie

**Abstract**

We present an analysis of long-term trends in surface ozone (O<sub>3</sub>) across Ireland, with specific focus on the Mace Head atmospheric research station, representative of Northern hemispheric background atmospheric conditions. Surface O<sub>3</sub> dataset was characterised using advanced

21 trajectory analysis and seasonal decomposition, revealing distinct seasonal and spatial patterns.  
22 Findings show a significant rising trend in surface O<sub>3</sub> at Irish urban sites over the past two  
23 decades but without a similar trend at coastal sites. Highest O<sub>3</sub> levels and exceedances were  
24 observed at remote coastal sites, which are less susceptible to influence from local and easterly  
25 emissions but heavily influenced by transboundary pollution and stratospheric intrusion.

26 At Mace Head, springtime O<sub>3</sub> levels exhibit a declining trend, whereas wintertime levels show  
27 a rising trend. Focussing on the clean sector, the springtime decline remains significant, but  
28 without corresponding clean sector rising wintertime trends, implying the rising winter trends  
29 occur in response to declining local, United Kingdom (UK) and European emissions. Advanced  
30 modelling tools are used to quantify O<sub>3</sub> source contributions, elucidating key drivers behind  
31 the observed changes. Characteristic springtime O<sub>3</sub> maxima at Mace Head are predominantly  
32 attributed to stratospheric transport, hemispheric and long-range transport and lightning NO<sub>x</sub>.  
33 The complementary trend and sectoral observational analysis reveal a decline in total spring-  
34 time concentrations, with a more rapid decline in exceedances from the UK & continental  
35 sector.

36 This research highlights the importance of seasonal factors in air quality management across  
37 Ireland, emphasising the need for a multi-faceted approach to control O<sub>3</sub> levels and reduce  
38 exceedances through global and regional emission reductions.

39 **Keywords** – Surface ozone, Meteorology, NO<sub>x</sub>, Climate, CH<sub>4</sub>, Emissions, VOC.

## 40 **1. Introduction**

41 Surface Ozone (O<sub>3</sub>) has significant implications for health, vegetation, and climate. Its  
42 chemical production is driven by complex photochemical processes, responding non-linearly  
43 to pollution control, creating challenges for its effective regulation. Elevated O<sub>3</sub> levels cause  
44 severe health issues with prolonged exposure to high O<sub>3</sub> levels, causing respiratory issues,

45 cardiovascular problems, and reduced lung function, particularly in sensitive populations  
46 such as children, the elderly, and individuals with pre-existing respiratory conditions (Lin et  
47 al., 2018; Todorović et al., 2019; Zhang et al., 2019, WHO, 2021). O<sub>3</sub> pollution can adversely  
48 impact vegetation by reducing agricultural productivity (Ashmore et al., 2005; Paoletti et al.,  
49 2006). O<sub>3</sub> is also the third most significant greenhouse gas after Carbon Dioxide (CO<sub>2</sub>) and  
50 methane (CH<sub>4</sub>), contributing to climate instability (IPCC, 2021). O<sub>3</sub> can live for several  
51 weeks in the free troposphere, and it is affected by large-scale atmospheric circulation  
52 patterns (Wespes et al., 2017). Meteorological factors such as temperature, solar radiation,  
53 wind speed, and atmospheric stability play a significant role in O<sub>3</sub> formation. (Ding et al.,  
54 2023; Khiem et al., 2010).

55 O<sub>3</sub> is formed in the atmosphere from precursors Nitrogen Oxides (NO<sub>x</sub>), carbon monoxide  
56 (CO), and volatile organic compounds (VOCs) through photochemical reactions.  
57 Photochemical production requires solar radiation, and reaction rates are correlated with  
58 temperature, it is noted that little photochemical production occurs at temperatures below  
59 20°C (Coates et al., 2016). The reactive, interdependent atmospheric chemistry leads to a  
60 non-linear relationship between O<sub>3</sub> and its precursors (Seinfeld and Pandis, 2016), and  
61 effective O<sub>3</sub> mitigation requires an understanding of processes influencing O<sub>3</sub> production and  
62 removal mechanisms (Fowler et al., 2013).

63 NO<sub>x</sub> can suppress or enhance O<sub>3</sub> formation, depending on local atmospheric chemistry  
64 regime. In polluted urban environments, high NO<sub>x</sub> emissions can lead to O<sub>3</sub> dissociation,  
65 retarding photochemical O<sub>3</sub> formation (NO<sub>x</sub>-saturated/VOC-limited regime), with local  
66 pollution events potentially titrating surface O<sub>3</sub> completely, converting NO to NO<sub>2</sub>. This  
67 effect is more prevalent in wintertime, when temperatures are low, and there is little solar  
68 radiation to facilitate photochemical production of O<sub>3</sub>. In relatively clean environments, O<sub>3</sub>  
69 formation is correlated with NO<sub>x</sub> concentration (Tavella & da Silva Júnior, 2021).

70 O<sub>3</sub> and precursors can be transported over great distances in the troposphere, with transport  
71 from distant polluted regions accounting for 40% of O<sub>3</sub> abundance in remote regions (Sudo and  
72 Akimoto, 2007). While O<sub>3</sub> is transported in air masses from distant sources, although titration  
73 occurs, O<sub>3</sub> is replenished via mixing. Over Ireland, advection of O<sub>3</sub> rich air from continental  
74 outflow is a significant source of Irish O<sub>3</sub>, but European continental air-masses can also  
75 transport pollutants to trigger O<sub>3</sub> depletion events, but the O<sub>3</sub> depleting effect of the air masses  
76 originating from Europe is in decline in response to European Union (EU) pollution control  
77 strategies (Derwent et al., 2024).

78 Seasonal and regional variations further complicate the regulation, with higher O<sub>3</sub> levels  
79 observed in summer across the northern hemisphere due to increased temperatures, solar  
80 radiation, and abundant precursors (Moiseenko et al., 2021; Sicard et al., 2016). In marine  
81 boundary layers, O<sub>3</sub> levels are generally lower than in continental regions, though specific  
82 oceanic environments can exhibit high O<sub>3</sub> concentrations due to inflows from polluted areas  
83 (Boylan et al., 2014; Girach et al., 2020). Another factor which influences O<sub>3</sub> levels is the  
84 North Atlantic Oscillation (NAO), which influences O<sub>3</sub> levels in Western Europe. During a  
85 positive NAO phase, O<sub>3</sub> levels increase. In contrast, during a negative NAO phase (NAO-  
86 low), O<sub>3</sub> levels decrease. This effect is particularly notable in southwest, central, and northern  
87 Europe (Bonaccorso et al., 2015; Creilson et al., 2003; Pausata et al., 2012).

88 While Ireland's air quality is mostly governed by the influx of clean maritime air from the  
89 Atlantic Ocean (Tripathi et al., 2010), certain synoptic scenarios allow for the intrusion of  
90 polluted air masses from continental Europe. These events, though infrequent, can bring  
91 substantial amounts of O<sub>3</sub> and its precursors (NO<sub>x</sub> and VOCs), contributing to short-term O<sub>3</sub>  
92 pollution episodes.

93 The World Health Organisation (WHO) publishes Air Quality Guidelines (AQGs) as a non-  
94 legally binding global target for governments to achieve within their jurisdictions. These

95 AQGs comprise evidence-based recommendations of limit values to protect public health.  
96 The current recommended AQGs for O<sub>3</sub> is expressed as a daily maximum of 8-hourly running  
97 average O<sub>3</sub> value of 100 µg/m<sup>3</sup>. Days when O<sub>3</sub> levels exceed the recommended AQGs are  
98 classified as exceedance days. Factors contributing to exceedances include high solar  
99 radiation, stagnant air masses, and local emissions and regional and transboundary transport  
100 of O<sub>3</sub> and precursors. Globally, over the past 150 years, there has been a 40% increase in O<sub>3</sub>  
101 levels owing to rising precursor emissions. (Archibald et al., 2020; Griffiths et al., 2021;  
102 Young et al., 2013). Despite European Union emission reduction policies, O<sub>3</sub> pollution  
103 remains a problem, with over 94% of those living in European cities exposed to O<sub>3</sub> levels  
104 exceeding the WHO AQGs in 2022 (EEA 2024, WHO 2021). Over 22,000 premature deaths  
105 in the EU were attributable to short-term exposure to O<sub>3</sub> in 2021 (Soares et al., 2023).

106 Long-term O<sub>3</sub> measurement data from the Mace Head research station in Ireland reveal a  
107 distinct seasonal pattern with peaks during spring and lows in summer. (Derwent, 1998;  
108 Derwent et al., 1994, 2018). Historical trends show increasing baseline O<sub>3</sub> levels in the 1980s  
109 and 1990s, stability in the 2000s, and a decline in the 2010s (Derwent et al., 2013; Derwent  
110 et al., 2018). Recent observational and modelling data have identified a broad O<sub>3</sub> maximum  
111 in spring and early summer, aligning with peak stratospheric transport (Ansari et al., 2024;  
112 Lin et al., 2012; Russo et al., 2023). O<sub>3</sub> dynamics are complex, and studies reveal  
113 discrepancies between model output and observations (Bessagnet et al., 2016; Vautard et al.,  
114 2012), highlighting the need for further understanding of factors governing O<sub>3</sub> levels and  
115 trends.

116 This study investigates the distribution and trends of O<sub>3</sub> and its precursors across Ireland,  
117 providing valuable insights into the regional and hemispheric impact on Irish surface O<sub>3</sub> levels  
118 and exceedances. We analyse the long-term O<sub>3</sub> observational dataset for Ireland, identifying  
119 the mean and range of O<sub>3</sub> levels and seasonal variation at each site. We identify the long-term

120 annual and monthly trends at each site, quantifying the significance of each trend according to  
121 TOAR guidelines (Chang et al., 2023). We also identify the frequency and seasonality of  
122 exceedance of the WHO AQGs for the protection of human health from O<sub>3</sub> pollution for each  
123 site. The trends and seasonality are contextualised by looking at trends in the Irish instrumental  
124 record of dominant precursors of NO<sub>x</sub> and CH<sub>4</sub> and we discuss the relationship between NO<sub>x</sub>  
125 and O<sub>3</sub> by comparing anomalies between monthly average NO<sub>x</sub> and O<sub>3</sub> during lockdown  
126 compared to the average values for the three years prior to lockdown. Advanced modelling  
127 results using the Tropospheric Ozone Attribution of Sources with Tagging 1.0 (TOAST 1.0)  
128 framework (Butler et al., 2018; Butler et al., 2020) are validated against measurements at  
129 various sites for simulation period 2000-2018, and the simulation results analysed to determine  
130 the geographical and sectoral source of precursors that contribute to simulated O<sub>3</sub> at Mace  
131 Head, identifying seasonality and long-term trends in the sources. Finally, the observational  
132 data are classified using advanced trajectory clustering methods to separate air masses from  
133 the clean sector from those influenced by local, UK or EU emission sources, with seasonal  
134 trends identified for both clean and polluted sectors, and the exceedances classified as coming  
135 from either the clean or polluted sector.

136 This research highlights significant seasonal and temporal variations and long-term trends in  
137 O<sub>3</sub> concentrations, and the integrated approach, including observational and modelling  
138 analysis, enhances our understanding of the drivers of O<sub>3</sub> concentrations, trends and  
139 exceedances over Ireland and allows quantification of global and regional contributions to Irish  
140 surface O<sub>3</sub>.

141

## 142 **2. Data and methodology**

### 143 **2.1 Observational Network and Analysis Approach**

144 Measurement data is obtained from the Environmental Protection Agency, Ireland (EPA)  
145 (<https://eparesearch.epa.ie/safer/>). The O<sub>3</sub> monitoring network shown in Figure 1 has been  
146 operational in Ireland since 1994. O<sub>3</sub> is measured using an API M400 and O<sub>3</sub> analyser based  
147 on UV photometry at all monitoring sites. At the Mace Head site, O<sub>3</sub> was measured using a  
148 Monitor Labs 8810 analyzer and a Thermo Environmental O<sub>3</sub> monitor. Measurement operating  
149 accuracy is within 1.0 ppb, based on precision, calibration and drift characteristics.  
150 Measurements of O<sub>3</sub> precursors from EPA air quality monitoring sites are also monitored. The  
151 details of the measurements sites are shown in Table 1. Numerous previous studies have  
152 analysed this data, with a particular focus on the analysis of Mace Head data to assess  
153 background O<sub>3</sub> (Carslaw, 2005; Derwent et al., 1994, 1998, 2001, 2004, 2008, 2013; Derwent,  
154 Manning, Simmonds, & Doherty, 2018; Derwent, Manning, Simmonds, Spain, et al., 2018;  
155 Oltmans et al., 2013; Simmonds et al., 2004; Tripathi et al., 2010, 2012, 2013). CH<sub>4</sub> data is  
156 obtained from the Integrated Carbon Observation System (ICOS) network, accessible at  
157 [https://www.icoscp.eu/data-products/ATM\\_NRT\\_CO2\\_CH4](https://www.icoscp.eu/data-products/ATM_NRT_CO2_CH4).

158 For this analysis, the observational sites were classified into three categories: Coastal, Rural,  
159 and Urban, as shown in Figure 1. The classification of the sites is based on Spohn et al., 2022,  
160 with the addition of the coastal category. Hourly data were used to evaluate annual trends  
161 based on monthly mean concentrations. Seasonal analysis conducted for the four main  
162 meteorological seasons in Ireland, namely Spring (March, April, May), Summer (June, July,  
163 August), Autumn (September, October, November), and Winter (December, January,  
164 February). O<sub>3</sub> exceedances were calculated based on the WHO AQGs, indicating that the  
165 maximum daily average over eight hours (MDA8) should not exceed 100 µg/m<sup>3</sup>. A  
166 significant analysis was performed on data measured at the Mace Head Atmospheric

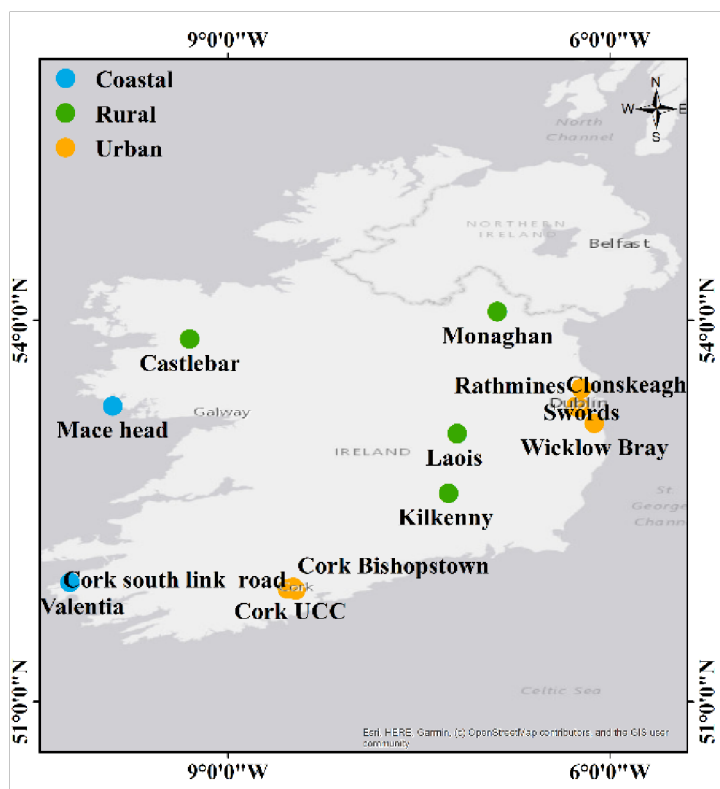
167 Research Station (53°33'N, 9°54' W), which is exposed to pristine marine air masses  
 168 approximately half of the time. (Grigas et al., 2017; O'Dowd et al., 2014).

169 **Table 1.** Details of Environmental Protection Agency Ireland (EPA) O<sub>3</sub> measurement sites  
 170 over Ireland, with location information, and the data period used for the study.

171

Site	Data availability	Type	Latitude	Longitude
Mace head	1994-2022	Coastal	53.3253	-9.9036
Valentia	2001-2022	Coastal	51.9385	-10.24
Monaghan	1995-2022	Rural	54.0661	-6.883
Laois	2005-2022	Rural	53.1076	-7.1983
Kilkenny	2012-2022	Rural	52.6383	-7.2676
Castlebar	2009-2022	Rural	53.851	-9.3003
Rathmines	2002-2022	Urban	53.322	-6.2672
Clonskeagh	2008-2022	Urban	53.3118	-6.2353
Swords	2009-2022	Urban	53.4631	-6.2222
Wicklow Bray	2009-2022	Urban	53.1873	-6.122
Cork South link road	2014-2022	Urban	51.8785	-8.4649
Cork Bishops town	2016-2022	Urban	51.8858	-8.53321
Cork UCC	2018-2022	Urban	51.9	-8.4863

172



173

174

175 **Figure 1.** – The map of EPA O<sub>3</sub> measurement sites over Ireland with classification of  
 176 backgrounds.

177 Trend analysis was conducted using the Openair package in R, software designed for the  
 178 analysis of atmospheric composition data. Trends were determined using the Theil-Sen slope  
 179 estimator and Mann-Kendall tests to quantify significance, in accordance with the  
 180 Tropospheric Ozone Assessment Report (TOAR) guidelines (Lefohn et al., 2018). Theil-Sen  
 181 directly provides a robust slope estimate. It is a reliable measure of change over time for  
 182 direct interpretation and comparison. It is a robust method for estimating trend slopes in time  
 183 series data, preferable to traditional least-squares regression, which can be sensitive to  
 184 extreme values and outliers. Uncertainty or reliability of the trend is calibrated according to  
 185 the p-value, as outlined by (Chang et al., 2023), consistent with the best statistical practices  
 186 for analysis used in the second phase of TOAR.

## 187 **2.2 Clean air sector identification from Back trajectories**

188 Baseline O<sub>3</sub> refers to the concentration of O<sub>3</sub> in air masses minimally influenced by local or  
189 regional anthropogenic emissions. Back-trajectory methods are widely used to estimate  
190 baseline O<sub>3</sub> levels by analysing the origins and transport pathways of air masses reaching  
191 observation sites. Typically, Lagrangian dispersion models are used to trace air parcels  
192 backwards in time and identify their origin.

193 For this study, air mass trajectories arriving at Mace Head were calculated using the Hybrid  
194 Single Particle Lagrangian Integrated Trajectory Model (HYSPLIT) (Draxler et al., 2003; Stein  
195 et al., 2015) in conjunction with R software. The air masses were classified into two categories:  
196 the clean sector and EU-influenced sector. An air mass was considered part of the clean sector  
197 when air mass trajectories remained over the ocean surface for the previous 72 hours, the  
198 remaining air mass trajectories are classified as the EU-influenced sector. Meteorological data  
199 for the analysis were derived from NOAA reanalysis data (Stunder et al., 2004). The 72h  
200 duration captures regional/long-range transport without trajectory error from meteorological  
201 uncertainties. The 100 m height was used to represent the well-mixed flow of the boundary  
202 layer above the surface. The 06:00 UTC aligns with synoptic times and can match daily O<sub>3</sub>  
203 cycles or measurement periods. Therefore, calculations were performed for 6:00 UTC each  
204 day, with a final trajectory height of 100 meters, covering the years 2000 to 2022. The O<sub>3</sub>  
205 concentrations observed during the clean sector were averaged to derive baseline levels,  
206 consistent with previous studies on baseline O<sub>3</sub> trends and sources (Derwent et al., 2013;  
207 Oltmans et al., 2006).

## 208 **2.3 CAM4-Chem Model**

209 The CAM-Chem air quality model, part of the Community Earth System Model (CESM),  
210 simulates atmospheric chemistry and the interactions among chemical constituents,  
211 meteorology, and climate. It incorporates detailed chemical mechanisms, emission inventories,

212 and meteorological data to simulate pollutant dispersion, thereby allowing us to determine air  
213 quality trends. CAM-Chem has been applied in numerous studies, significantly contributing to  
214 the understanding of regional and global atmospheric processes. (Lamarque et al., 2012; Tilmes  
215 et al., 2016). The model features a flexible chemical pre-processor to allow for detailed  
216 handling of atmospheric chemistry. Studies have demonstrated that CAM-Chem accurately  
217 represents conditions in both the troposphere. (Aghedo et al., 2011; Lamarque et al., 2010) and  
218 the stratosphere (Lamarque et al., 2008; Lamarque and Solomon, 2010), including temperature  
219 structure and dynamics (Butchart et al., 2011). Offline CAM-Chem has also been utilised in  
220 the Hemispheric Transport of Air Pollution (HTAP) assessments. (Anenberg et al., 2009; Fiore  
221 et al., 2009; Jonson et al., 2010; Shindell et al., 2008; Tan et al., 2018).

222 For the current study, we analyse simulations of the Community Atmospheric Model version 4  
223 CAM4-Chem (Community Atmosphere Model version 4 with chemistry) (Lamarque et al.,  
224 2012). The model simulations were carried out at a horizontal resolution of  $210 \text{ km} \times 280 \text{ km}$ ,  
225 with 56 vertical levels for the 2000-2018 period, with specified dynamics derived from  
226 MERRA2 reanalysis. (Molod et al., 2015). The coarse spatial resolution, it is still appropriate  
227 because it provides a reliable representation of the regional background atmosphere  
228 influencing the urban sites. The model captures large-scale features of atmospheric transport,  
229 seasonal variability, and background  $\text{O}_3$  levels, all of which are essential for interpreting urban  
230 observations. Comparing urban measurements with regional-scale CAM-Chem outputs allows  
231 local pollution effects to be distinguished from regional atmospheric influences.

232 Tagged source attribution of tropospheric ozone (TOAST 1.0) is a novel tagging methodology  
233 developed for the CESM to quantify source contributions to  $\text{O}_3$ . Unlike traditional methods  
234 that rely on sensitivity simulations, TOAST uses an online tagging approach to track  $\text{O}_3$   
235 production from specific  $\text{NO}_x$  and VOC sources (e.g., anthropogenic, biogenic, biomass  
236 burning, lightning) directly within the model, allowing for efficient attribution of  $\text{O}_3$  to regional

237 and sectoral emissions while maintaining full chemical coupling. The tool has been validated  
238 against observations and demonstrates utility in disentangling the impacts of different emission  
239 sectors on O<sub>3</sub> pollution. (Butler et al., 2018, 2020; Lupaşcu et al., 2022; Nalam et al., 2025).

240 The model results have inherent standard uncertainties common in any modelling exercise, i.e.,  
241 uncertainties in emission inventories in terms of magnitude and spatial accuracy; uncertainties  
242 in model parameters, e.g., surface resistance for deposition for various surfaces, boundary layer  
243 mixing, photolysis, chemical kinetic parameters (Wild et al., 2025) and structural deficiencies  
244 such as a coarse resolution and missing processes (e.g., halogen chemistry; Saiz-Lopez et al.,  
245 2025). The sum of all tagged contributions very closely matches the total simulated ozone  
246 (which is simulated independently and is not just an algebraic sum of the tagged contributions)  
247 is a good validation of the tagging mechanism (Butler et al., 2018 and 2020).

248 Global CAM4-Chem model simulations are performed for the years 2000-2018 with NO<sub>x</sub> and  
249 VOC tagging (as described in Ansari et al., 2025; Nalam et al., 2025), with the base chemical  
250 mechanism (Emmons et al., 2012) and source code modified to account for extra tagged species  
251 representing regional and sectoral identities. Anthropogenic emissions of NO<sub>x</sub>, CO, and non-  
252 methane volatile organic compounds (NMVOCs) are incorporated from the Hemispheric  
253 Transport of Air Pollution version 3 emissions inventory. (HTAPv3; Crippa et al., 2024), which  
254 includes land-based emissions, international shipping emissions, and aircraft emissions.  
255 Biomass burning emissions are sourced from the GFED-v4 inventory (Van Der Werf et al.,  
256 2010), while biogenic NMVOC emissions are derived from CAM4-GLOB-BIOv3.0. The O<sub>3</sub>  
257 source attribution technique used for this study is described in (Butler et al., 2020).

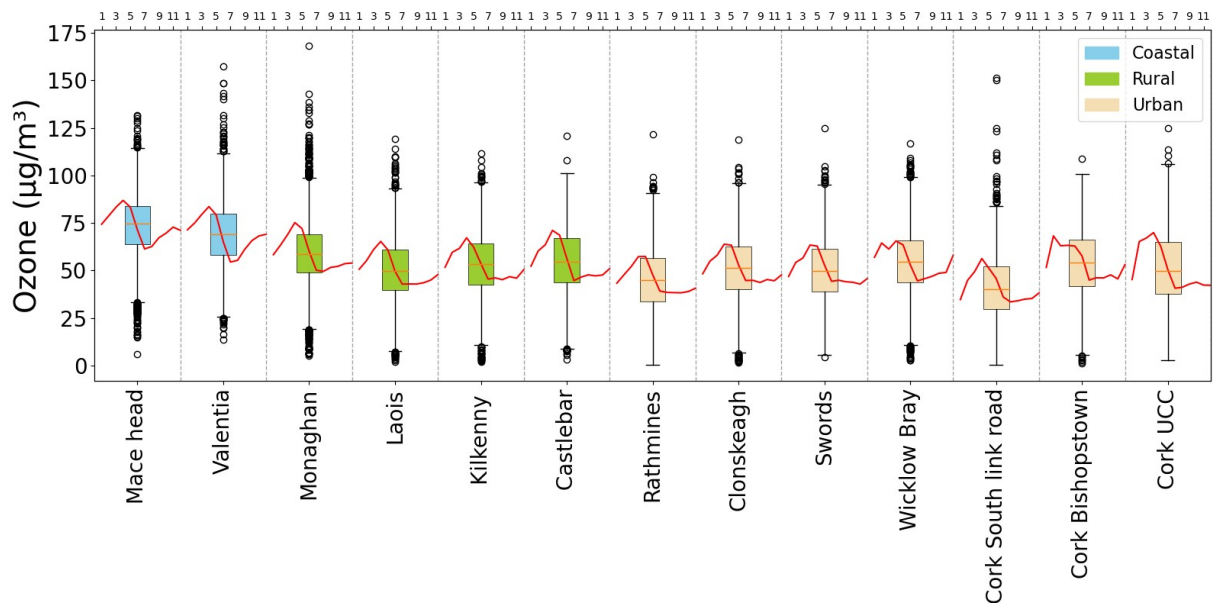
258

## 259 **3. Results and discussions**

### 260 **3.1 Yearly variation of O<sub>3</sub>**

261 Figure 2 shows box plots, illustrating the average O<sub>3</sub> concentrations for 13 sites over the  
262 duration of the available dataset, as discussed in section 2.1, providing a comprehensive  
263 overview of the variability and distribution of O<sub>3</sub> concentration. Coastal sites, Mace Head and  
264 Valentia, show higher O<sub>3</sub> levels compared to other sites, with annual average concentrations of  
265 77 µg/m<sup>3</sup> and 69 µg/m<sup>3</sup>, respectively. In urban areas like Rathmines, Dublin, O<sub>3</sub> concentrations  
266 remained consistently lower, with averages ranging from 39 to 56 µg/m<sup>3</sup>. Similarly, South link  
267 road and Bishopstown sites in Cork city, recorded relatively lower concentrations compared to  
268 coastal and rural locations, reflecting the impact of high urban NO<sub>x</sub> emissions. Rural sites like  
269 Laois and Kilkenny showed intermediate O<sub>3</sub> concentrations, less influenced by urban  
270 emissions. These sites consistently show O<sub>3</sub> averages ranging between 50 to 57 µg/m<sup>3</sup>, with  
271 little variability, highlighting the predominant role of steady background O<sub>3</sub> contributions in  
272 rural sites. O<sub>3</sub> concentrations vary significantly with proximity to emission sources adjacent  
273 to urban areas. O<sub>3</sub> levels can be lower due to titration, where O<sub>3</sub> reacts with NO, causing O<sub>3</sub>  
274 depletion, but the transport of precursors can cause an increase in O<sub>3</sub> concentration downwind  
275 of the sources (Jeon et al., 2014; Monks et al., 2015; Zhu et al., 2012)

276 The red line over the box shows a clear seasonal pattern in O<sub>3</sub> concentration for each site,  
277 with a springtime (March-April) peak and summertime (June-July) dip, with the highest  
278 peaks in the coastal sites, and lowest dips in urban sites, influenced by local emissions, e.g.  
279 Cork South Link Road and Swords.



280

281 **Figure 2.** Annual average O<sub>3</sub> concentration at different sites in Ireland. In each box, the  
 282 lowest whisker level represents the 5<sup>th</sup> percentile, the box spans from the 25<sup>th</sup> to the 75<sup>th</sup>  
 283 percentile, the horizontal line within the box represents the median 50<sup>th</sup> percentile, and the  
 284 upper whisker represents the 95<sup>th</sup> percentile. The average of monthly O<sub>3</sub> values calculated  
 285 for the entire period of each station, and the red line shows the average monthly O<sub>3</sub> variation  
 286 of all sites top axis shows the month (1– 12).

287

### 288 3.2 O<sub>3</sub> Trend analysis

#### 289 3.2.1 Yearly trend

290 Table 2 summarises the Theil-Sen trends in O<sub>3</sub> concentration (in µg/m<sup>3</sup> per year) across 13  
 291 monitoring sites in Ireland over different periods: 5 years (2018-2022), 10 years (2013-2022),  
 292 15 years (2008-2022), and the available years of data for each site. In the coastal regions,  
 293 Mace Head shows a consistent decrease in O<sub>3</sub> levels over the 5, 10, and 15-year periods,  
 294 although the entire dataset exhibits a small rising trend 0.02 µg/m<sup>3</sup> per year. These trends are  
 295 mostly in agreement with previous studies, where there was a positive trend observed in

326 background O<sub>3</sub> up to the mid-2000s, which stabilised and began to decline in the 2010s  
327 (Derwent et al., 2018). Valentia shows a long-term decreasing trend of -0.23 µg/m<sup>3</sup> per year,  
328 consistent with the previous study by Tripathi et al..2010.

329 In rural areas, Monaghan exhibits a declining trend in O<sub>3</sub> concentrations across all time  
330 periods, indicating an overall reduction. Laois shows an upward trend over the 10 and 15-  
331 year periods, though there is a slight decline in the most recent 5 years. Kilkenny presents  
332 slight negative trends over the 5 and 10-year periods (-0.29 and -0.01µg/m<sup>3</sup> per year).  
333 Negative trends are observed in Castlebar-(0.71 and -0.05 µg/m<sup>3</sup> per year).

334 The Dublin urban area sites (Rathmines, Clonskeagh, Swords) predominantly show increasing  
335 trends in O<sub>3</sub> levels, indicative of changes in urban pollution or local emissions, with decreased  
336 suppression of O<sub>3</sub> levels in urban regions due to decreased local emissions. (Derwent et al.,  
337 2024). This is consistent with the “weekend effect,” as observed by Atkinson-Palombo et al.  
338 (2006) whereby a reduction in NO<sub>x</sub> due to reduced weekend traffic decreases O<sub>3</sub> removal by  
339 NO<sub>x</sub> titration, leading to higher surface O<sub>3</sub> levels, likely to occur in wintertime, and in regions  
340 with low photochemical production due to low insolation and temperatures, as observed in  
341 Ireland. It is duly noted that NO<sub>x</sub> driven O<sub>3</sub> removal dominates over photochemical production  
342 in these sites. A comparable study in the UK carried out by Finch and Palmer (2020) attributed  
343 similarly rising trends in surface O<sub>3</sub> between 1999 and 2019 to decreasing NO<sub>x</sub>, characterising  
344 UK observation sites as VOC-limited. Mixed results are observed at the urban stations of Cork,  
345 with a positive trend at the South-link Road and Bishopstown, but a negative trend at the UCC  
346 station. The UCC station, although classified as urban, lies within the university campus;  
347 hence, it would not be subject to significant local emissions. Coastal sites like Mace Head and  
348 Valentia generally show decreasing trends, potentially due to less local emission sources but  
349 with more significant impacts from regional and long-range transport, However, a detailed  
350 analysis of the trends requires consideration of seasonal effects.

321 Previous studies indicate that in northeast Europe, peak surface ozone concentrations have  
 322 generally declined, reflecting the effectiveness of emission control measures (Yan et al., 2018).  
 323 In contrast, background and lower-level O<sub>3</sub> concentrations have continued to increase,  
 324 particularly at rural and suburban sites. Consistent with this, urban observations from 2000–  
 325 2021 show increasing trends in median and lower-percentile O<sub>3</sub> levels, while the highest  
 326 extremes have mostly decreased (Nelson et al., 2025).

327 **Table 2.** - Trends in surface O<sub>3</sub> concentration (µg/m<sup>3</sup> per year) calculated for 13 sites in Ireland  
 328 over different periods over the complete dataset: 5 years (2018-2022), 10 years (2013-2022),  
 329 15 years (2008-2022), and the available measurement record for the site. The p-value evaluates  
 330 the reliability of the trend, whereas a lower p-value indicates trend certainty. Adopting the trend  
 331 reliability scale defined for TOAR-II studies (Chang et al., 2023), trends with very high  
 332 certainty will be marked by \*\*\*( $p \leq 0.001$ ), trends with high certainty with \*\* ( $p \leq 0.01$ ), and  
 333 low to medium certainty with \*( $p \leq 0.05$ ).

334  
 335  
 336

Site No	Site name (Classification)	Measurement Record	Trend over record µg/m <sup>3</sup> per year	5-year trend 2018-2022 µg/m <sup>3</sup> per year	10-year trend 2013-2022 µg/m <sup>3</sup> per year	15-year trend 2008-2022 µg/m <sup>3</sup> per year
1	Mace Head (C)	1994-2022	0.02	-0.25	-0.31***	-0.11*
2	Valentia (C)	2001-2022	-0.23****	-1.15****	-0.84****	-0.32***
3	Monaghan (R)	1995-2022	-0.19****	-0.74**	-0.35**	-0.09*
4	Laois (R)	2005-2022	0.39****	-0.15	0.3**	0.46****
5	Kilkenny (R)	2012-2022	0.02	-0.29	-0.01	

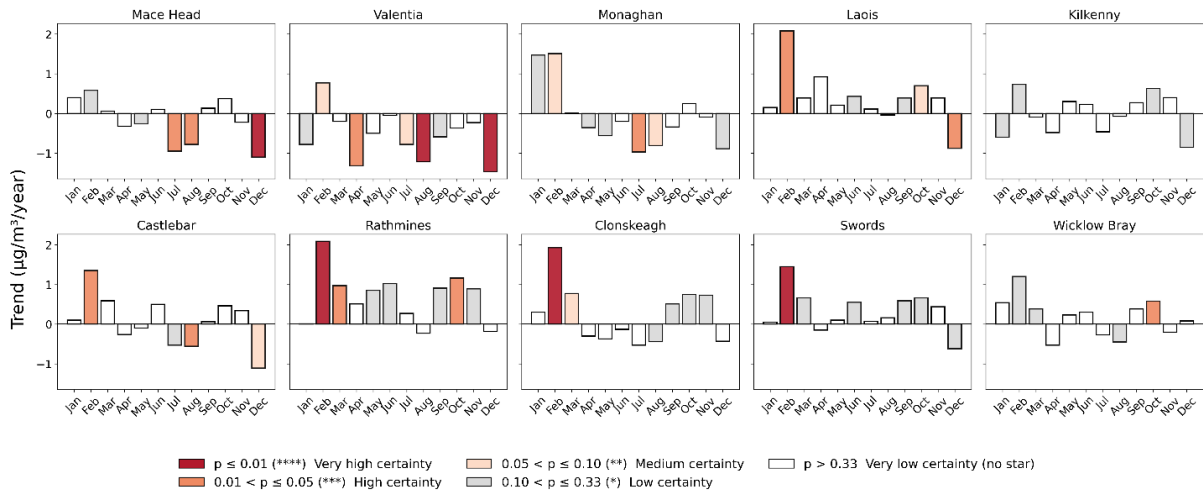
6	Castlebar (R)	2009-2022	0.18**	-0.71**	-0.05	
7	Rathmines (U)	2002-2022	0.27*****	1.72*****	1.15*****	0.48*****
8	Clonskeagh (U)	2008-2022	0.33*****	0.97***	0.12	0.33*****
9	Swords (U)	2009-2022	0.6*****	0.07	0.33*****	
10	Wicklow Bray (U)	2009-2022	0.14*	0.04		
11	Cork Southlink Road (U)	2014-2022	0.51*	-0.44		
12	Cork Bishopstown (U)	2016-2022	1.05***	-1.81*		
13	Cork UCC (U)	2018-2022	-0.94	-0.94		

337

338

### 339 3.2.2 Monthly trend

340 Figure 3. shows the monthly trend for 10 years from the period 2012-2022. Mace Head  
341 (coastal) and Monaghan (rural) sites predominantly show a rising trend in winter/early spring,  
342 with a decreasing trend in late spring to summer. Valentia shows a decreasing trend in every  
343 month except February when levels are significantly impacted by long-range transport and  
344 stratospheric sources (Auvray and Bey, 2005; Pan et al., 2018). Urban sites show a general  
345 increasing monthly trend, as yearly trend but with a seasonal signal in Clonskeagh with an  
346 increase in winter-spring and a decrease in late spring or summer. Laois is characterised as a  
347 rural site yet exhibits rising trends similar to the urban sites for all months except December,  
348 indicating that the measurement station is affected by nearby emissions. Seasonal trends of the  
349 15-year dataset are supplied in supplementary Figure S2, where coastal stations exhibit a  
350 pronounced increase in late winter, and a decrease in late summer, with a consistent near-year-  
351 round increase in Rathmines and Laois.



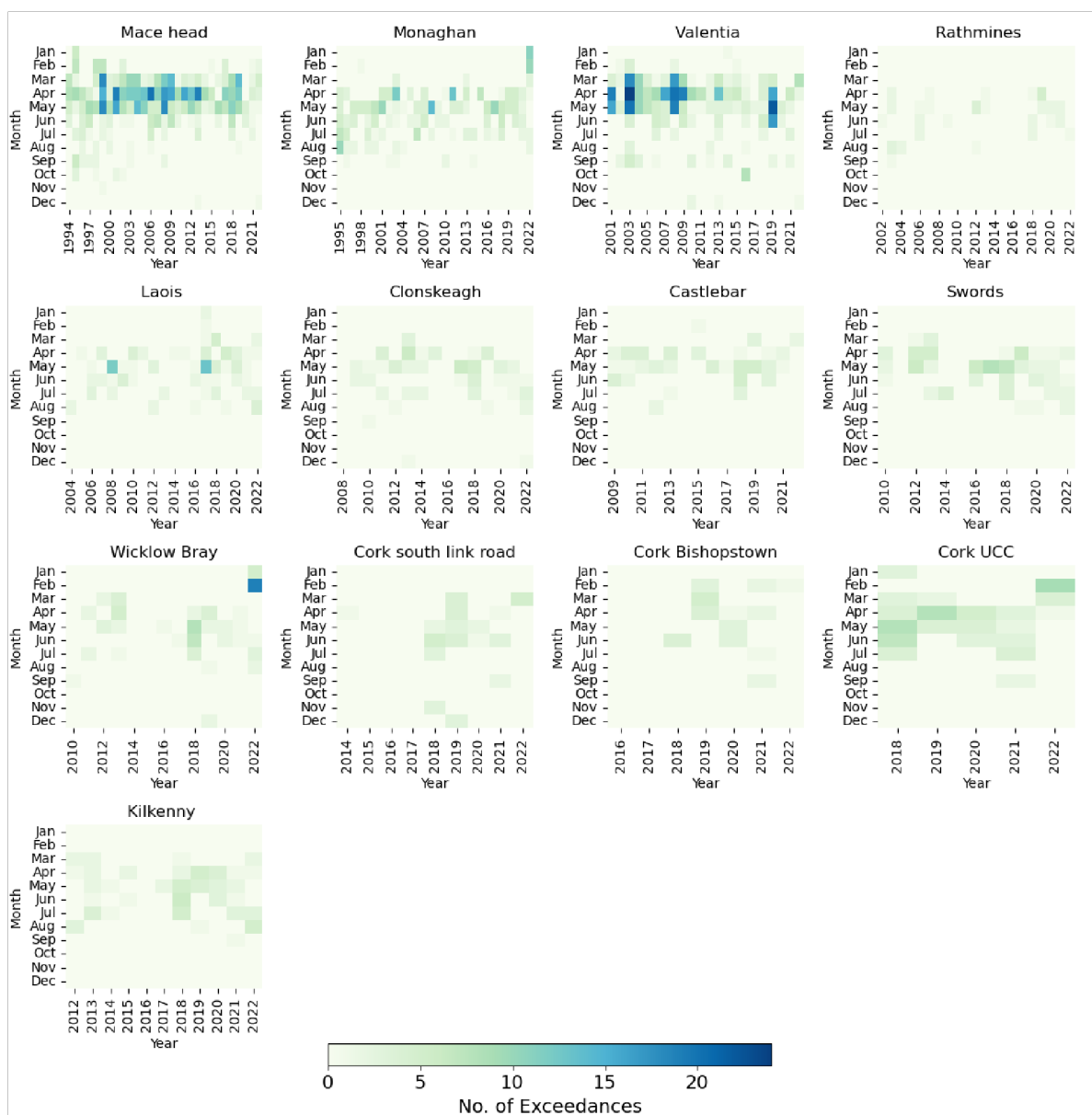
352

353 **Figure 3.** Monthly trend analysis of O<sub>3</sub> at different sites for 10 year period. (2012-2022)  
 354 Adopting the trend reliability scale defined for TOAR-II studies (Chang et al., 2023), trends  
 355 significance highlighted by colour.

356 **3.3 O<sub>3</sub> Exceedance**

357 The O<sub>3</sub> exceedances at 13 sites in Ireland over the available measurement dataset are identified  
 358 according to WHO criteria, and the results are shown in Figure 4. The highest and lowest  
 359 numbers of O<sub>3</sub> exceedances were observed at Mace Head and Rathmines, representing coastal  
 360 and urban sites, respectively. Most exceedances occurred in spring (March, April, May),  
 361 coinciding with the spring-time maximum. Rathmines had its highest number of exceedances  
 362 in April 2019, while Laois reached a peak of 13 exceedances in May 2017. Castlebar and  
 363 Swords show increased exceedance occurrences in spring and early summer, particularly  
 364 notable spikes occurring in 2010, 2013, 2016, and 2019. Conversely, Wicklow Bray exhibited  
 365 a different pattern, showing significant spikes in February and March 2022, alongside  
 366 occasional exceedances during March, April, and May, for example, in 2012 and 2018. Cork  
 367 South link Road also recorded exceedances, particularly in March and notably in June. Cork  
 368 Bishoptown shows exceedances, especially in February and March 2019, while Cork UCC  
 369 recorded exceedances in April and May 2019. Kilkenny consistently exhibited exceedances

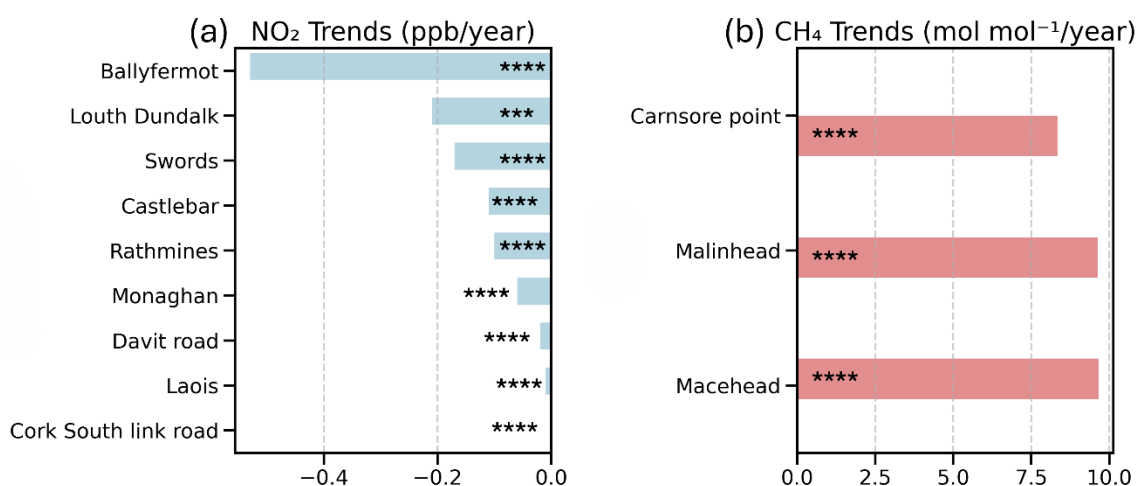
370 during spring and summer, with April and May often recording the highest numbers,  
 371 particularly in 2019. This highlights the impact of seasonal atmospheric conditions on O<sub>3</sub>  
 372 levels. It is noted that summertime exceedances, although less frequent in occurrence, indicate  
 373 photochemical production events that would be required to elevate O<sub>3</sub> levels from the annual  
 374 dip in the seasonal cycle to exceed the WHO AQG threshold. These episodic spikes are  
 375 characteristic of unique climatic or pollution events and warrant further study.



376  
 377 **Figure 4.** - Monthly O<sub>3</sub> exceedance at different sites in Ireland. Exceedances are defined as  
 378 days on which the maximum 8-hour running average of ozone (O<sub>3</sub>) exceeds 100 µg/m<sup>3</sup>.

379 Figure 5 depicts trends in NO<sub>2</sub> and CH<sub>4</sub> concentrations across various Irish measurement sites.  
 380 NO<sub>2</sub> trend calculations are based on site-specific data periods, Cork South link Road (2014–  
 381 2022), Ballyfermot (2003–2022), Davit Road (2018–2023), Rathmines (1995–2024), Swords  
 382 (2011–2025), Laois (2014–2026), Castlebar (2003–2027), Louth Dundalk (2019–2028), and  
 383 Monaghan (2001–2029). For CH<sub>4</sub>, the data cover the period 2010–2022. Most monitored sites  
 384 exhibit a decreasing trend in NO<sub>2</sub> concentrations, most likely in response to pollution control  
 385 on transportation, industrial activities, and energy production in the EU and North America  
 386 (Coleman et al., 2013; Donlon et al., 2024). In contrast, CH<sub>4</sub> levels observed at three sites -  
 387 Mace Head, Malin Head, and Carnsore Point indicate a significant and persistent rise in CH<sub>4</sub>  
 388 concentrations. Mace Head is known for its clean Atlantic air and Malin Head, situated at  
 389 Ireland’s northern tip near the UK border, offers a unique position to observe both clean marine  
 390 air and transboundary pollution, whereas Carnsore Point in the southeast captures air masses  
 391 from both the UK and mainland Europe (Spohn et al., 2022). These NO<sub>2</sub> and CH<sub>4</sub> trends reveal  
 392 a dual dynamic: while NO<sub>2</sub> levels are decreasing due to effective emission controls, CH<sub>4</sub> levels  
 393 relentlessly rise, highlighting the need for enhanced mitigation strategies targeting CH<sub>4</sub>.

394



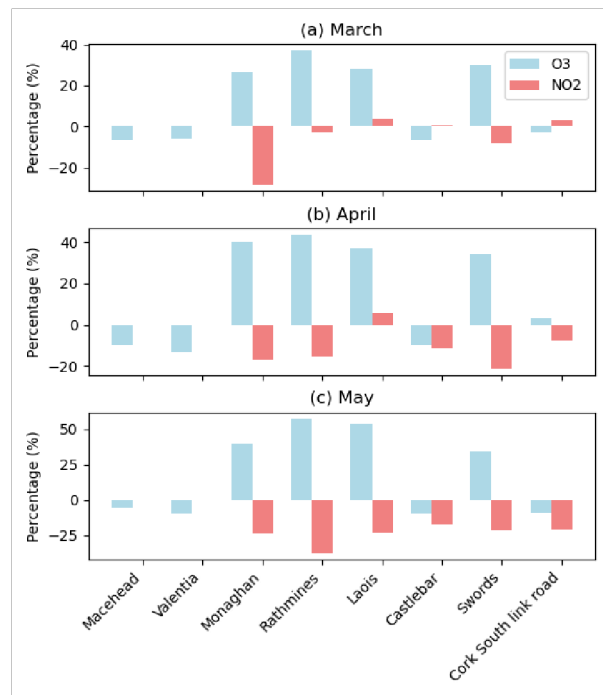
395

396 **Figure 5.** Trend in O<sub>3</sub> precursors NO<sub>2</sub> (a), and CH<sub>4</sub> (b) at different sites. Statistical significance  
397 of trend is indicated, increasing \* with increasing trend reliability, as in Table 1.

398 To evaluate the relationship between NO<sub>x</sub> and O<sub>3</sub> concentrations in an Irish context and the  
399 potential benefit of abrupt enforcement of NO<sub>x</sub> control measures, we assess the impact of the  
400 COVID-19 2020 lockdown, Spring 2020, whereby the lockdown period saw a prominent  
401 relative decrease in NO<sub>2</sub>, yet an increase in surface O<sub>3</sub> compared to average measurements for  
402 the same months 2017-2019 in most national monitoring stations (Figure 6). The negative  
403 correlation between O<sub>3</sub> and NO<sub>2</sub> is indicative of a NO<sub>x</sub>-saturated regime, normally associated  
404 with polluted urban environments and NO<sub>x</sub> titration events. Analysis of the impact of the  
405 lockdown on Irish O<sub>3</sub> is discussed by Spohn et al. (2022), and a similar O<sub>3</sub> decrease was widely  
406 observed across Europe during the COVID lockdown (Ordóñez et al., 2020; Tavella & da Silva  
407 Júnior, 2021; C. Zhang & Stevenson, 2022). Significant enhancement of O<sub>3</sub> occurs at the inland  
408 measurement sites, despite a 2020 springtime decrease in O<sub>3</sub> observed at background coastal  
409 sites, Mace Head and Valentia. These coastal stations are less sensitive to changes in European  
410 NO<sub>x</sub> emissions than inland sites and more sensitive to stratospheric and hemispheric transport  
411 (Tan et al., 2018). It is noted that April and May 2020 had unique meteorological conditions  
412 compared to previous years, with lower wind speed, less rain and significantly higher solar  
413 radiation, see Figure 12 in Spohn et al. (2022). These meteorological conditions would  
414 potentially facilitate photochemical O<sub>3</sub> production, contributing to positive O<sub>3</sub> anomalies  
415 during the lockdown period in addition to NO<sub>x</sub> reduction, also potentially enhancing dry  
416 deposition to the ocean. Further investigation into this topic would warrant model sensitivity  
417 studies, beyond the scope of this current work.

418 The negative correlation between NO<sub>x</sub> and O<sub>3</sub> under relatively clean atmospheric conditions  
419 indicates that O<sub>3</sub> levels are influenced predominantly by transport and chemical removal, and  
420 local photochemical production does not represent a significant surface O<sub>3</sub> source owing to

421 periods of low-insolation periods and low temperature, which are characteristic of Irish  
422 meteorology and frequent cloud cover (Pallé and Butler, 2002)



423

424

425 **Figure 6.** Percentage change in NO<sub>2</sub> and O<sub>3</sub> during the lockdown period of 2020 as compared  
426 to the 2017-2019 average at different sites in Ireland for (a) March (b) April (c) May.

### 427 3.4 Model and Observations Comparison

#### 428 3.4.1 Comparison between CAM4 – Chem Model and Observations

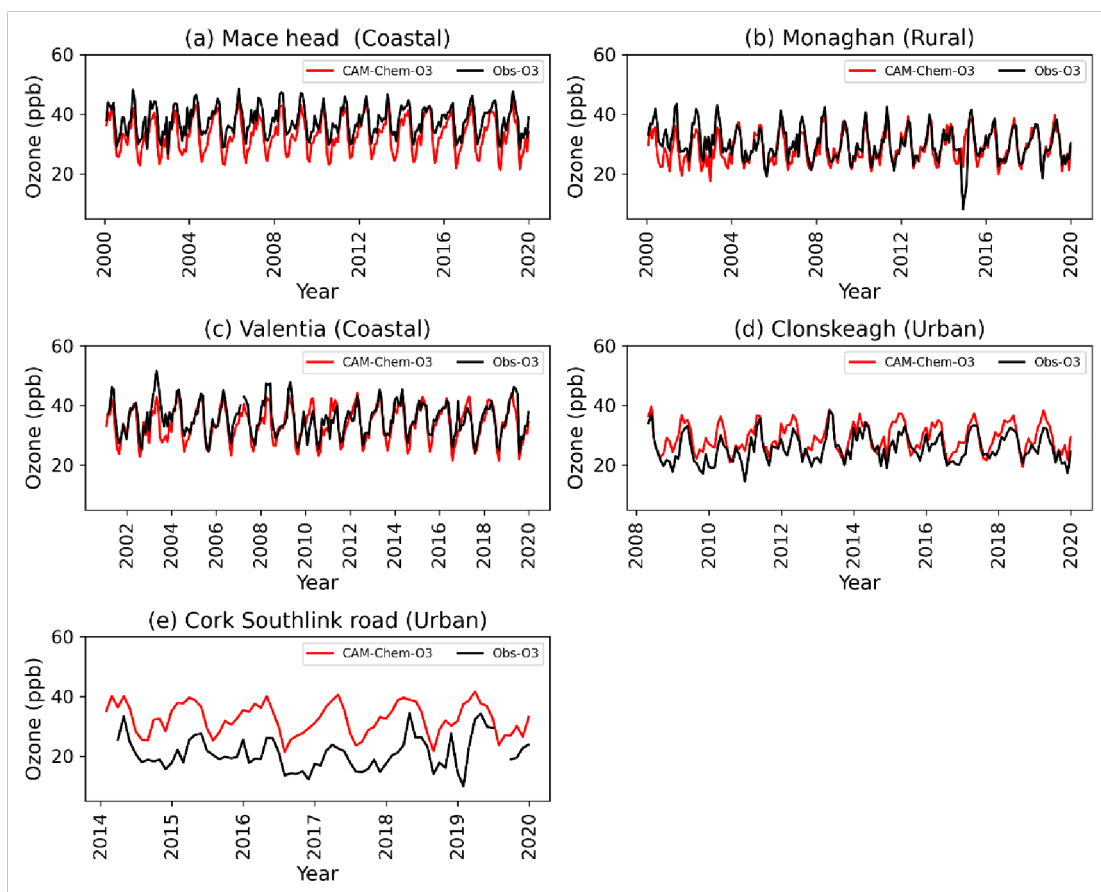
429 Global simulations were performed with the CAM4-Chem model enabled with source tagging  
430 (Butler et al., 2018) for 2000-2018 and the modelled O<sub>3</sub> over Ireland was compared with  
431 surface O<sub>3</sub> measurements at five sites. The grid details are shown in Figure S5. This model has  
432 already been comprehensively evaluated on a global scale (Nalam et al., 2025) against the  
433 TOAR-1 dataset, whereby the model performed well, albeit with a high bias attributable to the  
434 model's coarse resolution and implications for the resolution of urban chemistry (Ansari et al.,  
435 2025). Figure 7 shows the comparison of monthly O<sub>3</sub> CAM4-Chem and ground station O<sub>3</sub> data.

436 From this figure, it is observed that CAM4-Chem exhibits negative (positive) bias in rural and  
437 coastal (urban) sites. The underestimation at Mace Head is probably caused by the coarse grid  
438 resolution, covering a large area not representative of Mace Head conditions.

439 The influence of coastal meteorology, which suppresses O<sub>3</sub> formation due to cooler  
440 temperatures and persistent cloud cover (McVeigh et al., 2010), also leads to an  
441 underestimation of O<sub>3</sub> (Yerramilli et al, 2012). The dry deposition rate over land would exceed  
442 that over the ocean, leading to a lower simulated O<sub>3</sub> concentration for the grid cell which covers  
443 both land and ocean surface. Dry deposition is enhanced by solar radiation (Coleman et  
444 al.,2012; Coleman et al., 2013; Pio et al., 2000) hence model measurement discrepancy is at a  
445 maximum in late summer months. As outlined by Fiore et al. (2009), models average the  
446 landscape characteristics within a grid cell, which can enhance O<sub>3</sub> deposition and result in  
447 lower simulated O<sub>3</sub> concentrations; hence, the discrepancy is more pronounced in the clean  
448 sector data.

449 Overestimation of O<sub>3</sub> in Clonskeagh and Cork South link Road is likely due to coarse grid  
450 handling of localised emissions and subsequent atmospheric chemistry.

451



452

453 **Figure 7.** The comparison of Monthly CAM4 – Chem O<sub>3</sub> and Monthly O<sub>3</sub> observations at  
 454 five sites in Ireland.

455 At Mace Head, the model shows a negative mean bias of - 4.42 (-11.68% normalized mean  
 456 bias) but strong correlation ( $r=0.83$ ). In Monaghan and Valentia, the model shows smaller  
 457 biases of -1.43 and -1.54, normalized mean biases of -4.74% and -1.54%, and correlation  
 458 coefficients of 0.73 and 0.72, respectively. These high correlations are in line with (Tilmes  
 459 et al., 2015). However, at Clonskeagh and Cork South Link Road, the model overestimates,  
 460 with positive biases (3.38 and 11.98) and weaker correlations (0.68 and 0.49). Statistics are  
 461 given in more detail in the supplementary material Table S1. These results suggest better  
 462 model performance at coastal/rural sites and greater discrepancies in areas affected by local  
 463 sources, as expected at this model resolution.

464 **3.4.2 Source attribution using CAM4-Chem**

465 To quantify the contribution of various precursor emission sources to modelled O<sub>3</sub>  
 466 concentrations, the TOAST1.0 dual NO<sub>x</sub> and VOC tagging technique was utilised (Butler et  
 467 al., 2018). This allows attribution of modelled O<sub>3</sub> to the emissions of NO<sub>x</sub> and VOC precursors  
 468 across different source sectors and geographical regions as listed in Table 3.

469

470

471

472

473

474

475 **Table 3.** - List of tags used in NO<sub>x</sub> and VOC tagging.

<b>Regional Land-Based Tags</b>		<b>Regional Oceanic Tags</b>		<b>Global Sector/ProcessBased Tags</b>	
<b>ARC</b>	Arctic	<b>NAL</b>	North Atlantic	<b>AIR</b>	Aircraft
<b>CAS</b>	Central Asia	<b>ENA</b>	Eastern North Atlantic	<b>BIO</b>	Biogenic
<b>EAS</b>	East Asia	<b>NAE</b>	North America East Coast	<b>BMB</b>	Biomass Burning
<b>EUR</b>	Europe	<b>NAW</b>	North American West Coast	<b>LGT</b>	Lightning
<b>MCA</b>	Mexico & Central America	<b>NPA</b>	North Pacific	<b>STR</b>	Stratospheric Intrusion
<b>MDE</b>	Middle East	<b>BNS</b>	Baltic and North Seas	<b>XTR</b>	Extra untagged O <sub>3</sub>

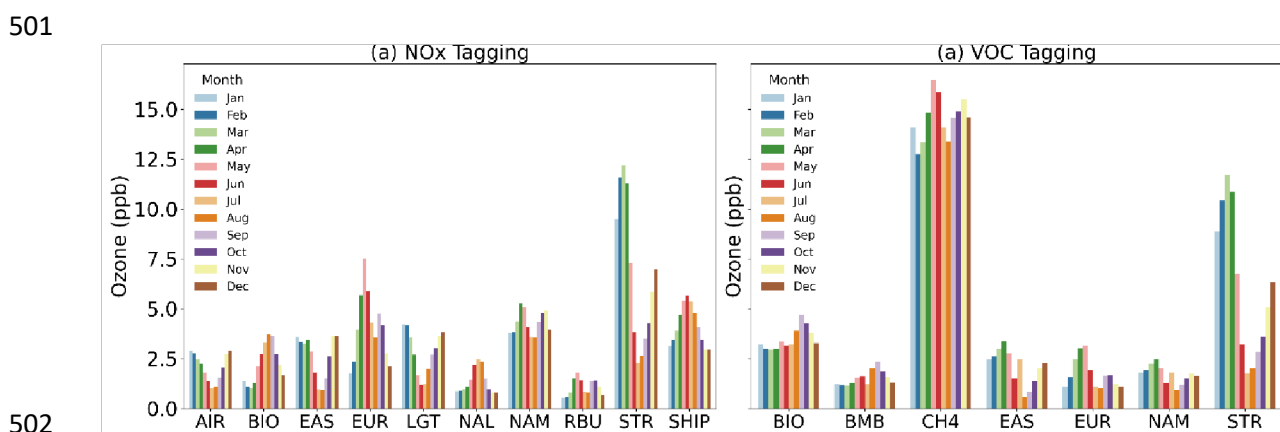
<b>NAF</b>	North Africa	<b>HBV</b>	Hudson Bay	<b>CH4</b>	Methane
<b>NAM</b>	North America	<b>IDO</b>	Indian Ocean	<b>OCN</b>	Oceanic Sources (DMS)
<b>RBU</b>	Russia-BelarusUkraine	<b>MBC</b>	Mediterranean, Black, and Caspian Seas	<b>SHV</b>	Shipping
<b>SAS</b>	South Asia	<b>SHO</b>	Southern hemispheric oceans	<b>AIR</b>	Aircraft
<b>SEA</b>	Southeast Asia			<b>INI</b>	InitialConditionO <sub>3</sub>
<b>VRW</b>	Rest of the World				

476  
477 The monthly tagged major precursor contributions to surface O<sub>3</sub> at Mace Head, averaged  
478 over the 2000-2018 simulation period, are shown in Figure 8. The stratospheric source of O<sub>3</sub>  
479 dominates in Winter-Spring, contributing to the spring-time maxima due to vigorous  
480 stratospheric transport. European NO<sub>x</sub> emissions contribution peaks in May, while lightning  
481 NO<sub>x</sub> has the greatest impact in winter. North American (NAM) NO<sub>x</sub> emissions contribute 3.5  
482 to 5.25 ppb, comparable to European NO<sub>x</sub>, but with an earlier peak in April. Aviation  
483 emissions contribute 1 to 3 ppb, with the highest contributions in winter and spring. Biogenic  
484 NO<sub>x</sub>, significant between June and October, contributes an average of 3.6 ppb, with higher  
485 contributions during August and September. O<sub>3</sub> derived from biogenic VOC sources average  
486 over 4 ppb during late autumn, maintaining a more sustained contribution throughout the  
487 year. East Asian NO<sub>x</sub> emissions, contribute up to 3.6 ppb, with a minimum contribution in  
488 July and August. North Atlantic shipping NO<sub>x</sub> (NAL) accounts for up to 2.4 ppb of O<sub>3</sub> during  
489 July. The total shipping NO<sub>x</sub> (SHIP) also contributes significantly It is the addition of all  
490 oceanic emissions and shows the highest contribution in June.

491 Methane (CH<sub>4</sub>) is the dominant reactive carbon molecule contributing to O<sub>3</sub> formation. VOC  
492 emissions from biomass burning also play a measurable role, contributing 1 to 2 ppb, with

493 their largest contributions in August and September. Finally, European VOC emissions  
 494 contribute 1 to 3 ppb, with the largest impact from March to May, coinciding with the spring-  
 495 time peak in surface O<sub>3</sub>.

496 These findings allow quantification of specific sources amidst the complex interplay of  
 497 regional and global sources in driving seasonal variations in surface O<sub>3</sub> levels over the Irish  
 498 domain, highlighting the roles of stratospheric processes, anthropogenic emissions, biogenic  
 499 sources, and lower-latitude contributions in shaping the observed patterns at background  
 500 monitoring sites such as Mace Head.



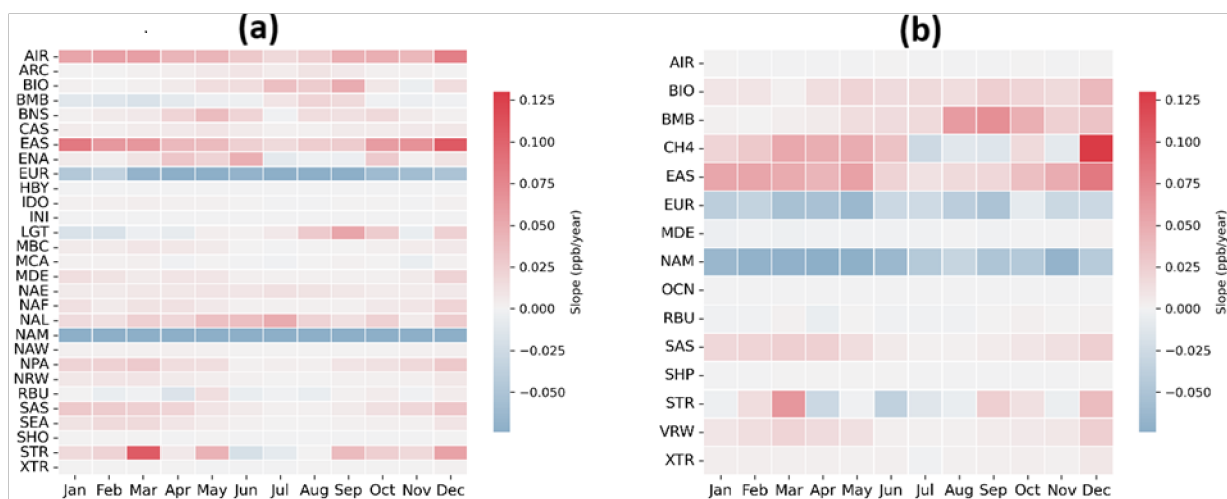
502  
 503 **Figure 8.-** Absolute contribution of major NO<sub>x</sub> sources (a) (NO<sub>x</sub> Tagging) and VOC source  
 504 (b) (NO<sub>x</sub> Tagging) to the CAM4-Chem simulated surface O<sub>3</sub> for the Mace Head grid cell  
 505 between 2000-2018.

506 Figure 9 shows the monthly changes in contributions to surface O<sub>3</sub> at Mace Head over the  
 507 simulation period (2000-2018). A negative (blue) trend indicates that the contribution of the  
 508 source to simulated surface O<sub>3</sub> in this grid cell has declined over the simulation period, whereas  
 509 a positive trend (red) indicates the contribution to surface O<sub>3</sub> has risen. Figure 9 (a) shows that  
 510 simulated O<sub>3</sub> at Mace Head originating from European or North American NO<sub>x</sub> decreases over  
 511 the simulation period, consistent with EU & North American emission reductions (Guerreiro

512 et al., 2014). More significant reduction occurs in late Spring through late summer, when EU  
513 NO<sub>x</sub> contributions are most significant to Mace Head O<sub>3</sub> concentrations (as seen in Figure 8).

514 There is a rising trend in simulated surface O<sub>3</sub> originating from NO<sub>x</sub> emissions from global  
515 aviation, from East Asia, and to a lesser extent from South Asia, which is more pronounced in  
516 the wintertime. Wintertime temperatures in South Asia are still high enough to sufficiently  
517 produce local ozone, especially when NO<sub>x</sub> emissions are rising (Crippa et al., 2023). The  
518 relatively longer, atmospheric lifetime of O<sub>3</sub> in the free troposphere during winter enables  
519 longer-range intercontinental transport (Huang et al., 2017; Yu et al., 2013). This seasonality in  
520 source contributions explains the observed reduction in spring-time maxima and increase in  
521 winter-time levels from the measurement record. EastAsian and South-Asian VOCs also  
522 contribute to a rising trend in simulated O<sub>3</sub>, with a more pronounced increase in winter and  
523 spring. This highlights a different pattern in hemispheric O<sub>3</sub> contributions, where EU and NAM  
524 emission reductions coincide with increased contribution from lower latitudes which could  
525 potentially become a more significant source of background O<sub>3</sub> in the Northern Hemisphere in  
526 the future. The contribution of CH<sub>4</sub> also has a positive trend over the simulation period, but the  
527 CH<sub>4</sub> trend has a reliable correlation only in December and spring periods, with very low  
528 certainty in CH<sub>4</sub> contribution trends for summer months (correlation coefficient,  $p > 0.33$ ).

529 Contribution to simulated O<sub>3</sub> at Mace Head from EU and NAM anthropogenic VOC show a  
530 negative trend for all months, consistent with trends from EU and NAM anthropogenic NO<sub>x</sub>.



531

532 **Figure 9.** Trends in contributions to monthly average modelled Mace Head grid cell surface

533 O<sub>3</sub> for the 2000-2018 period derived from (a) NO<sub>x</sub> tagging and (b) VOC tagging.

534 Table 4. shows the overall trend in the main contributors to NO<sub>x</sub> and VOC tagging. It is

535 observed that there is an increase in simulated surface O<sub>3</sub> originating from NO<sub>x</sub> contributions

536 from aviation and East Asia, while there is a decrease in European (EUR) and North American

537 (NAM) NO<sub>x</sub> contributions. In VOC tagging, Methane (CH<sub>4</sub>) and East Asian anthropogenic

538 VOC (EAS) contribute to a rising trend over the simulation period, whereas anthropogenic

539 VOC contributions from Europe (EUR) and North America (NAM) show a negative trend.

540 **Table 4** - Overall Trend in contributions to Mace Head grid cell O<sub>3</sub> simulated by CAM4-

541 Chem for NO<sub>x</sub> tagging and VOC tagging over the simulation period in units of ppb per year.

542 reliability scale defined for TOAR-II studies (Chang et al., 2023), Statistical significance of

543 trend is indicated using star notation: \*\*\*\* denotes  $p \leq 0.01$  (very high certainty), \*\*\* denotes

544  $0.01 < p \leq 0.05$  (high certainty), \*\* denotes  $0.05 < p \leq 0.10$  (medium certainty), \* denotes

545  $0.10 < p \leq 0.33$  (low certainty), and no star denotes  $p > 0.33$  (very low certainty).

NO <sub>x</sub> Tagging		VOC Tagging	
	Slope (ppb/year)		Slope (ppb/year)

AIR	0.0467****	CH <sub>4</sub>	0.0590****
EAS	0.0491****	EAS	0.0333****
EUR	-0.0900****	EUR	-0.0553****
NAM	-0.1243****	NAM	-0.0670****

546

### 547 3.5 O<sub>3</sub> Trends in Background and EU influenced sector Airmasses at Mace Head

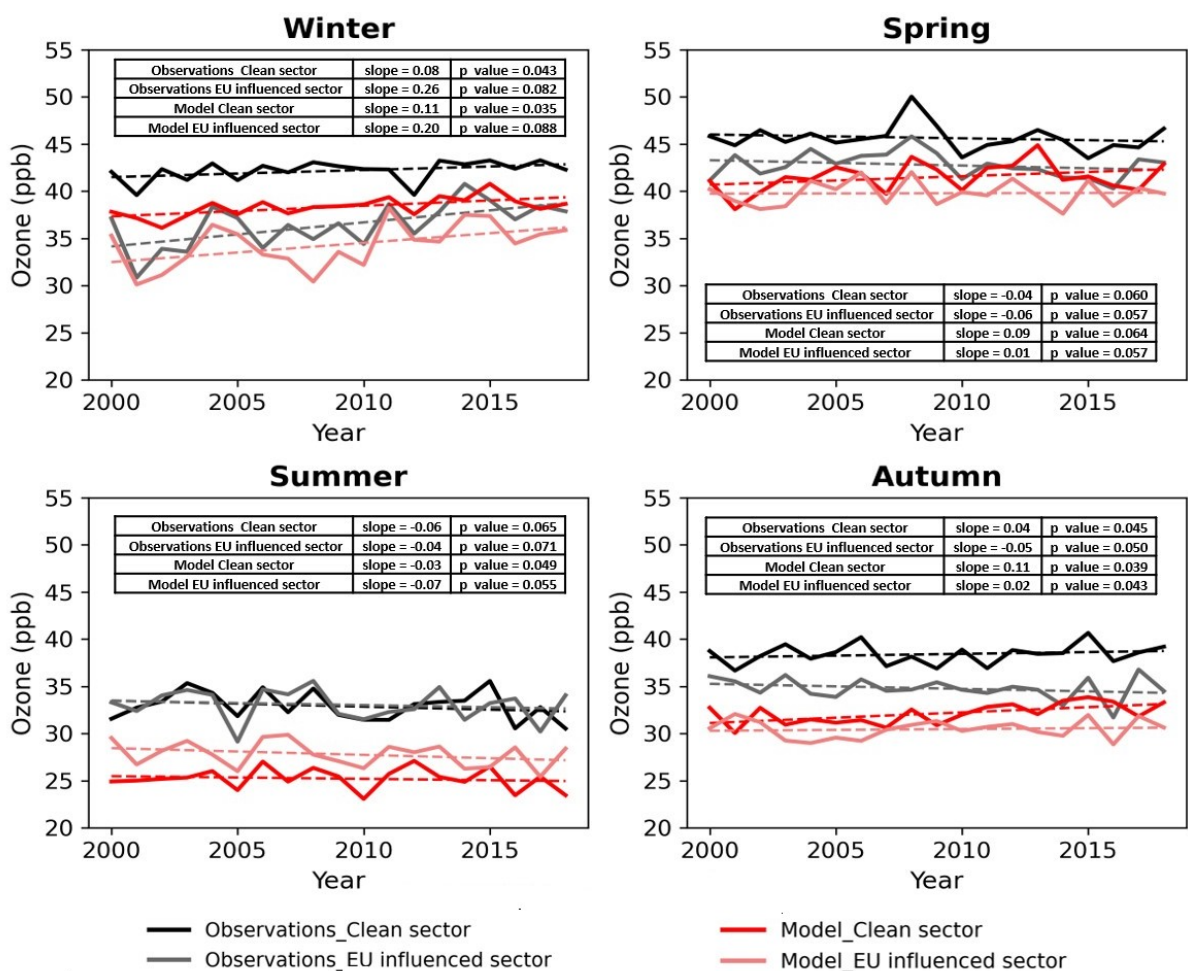
548 Although Mace Head is classified as a global background site, quantification of the baseline  
549 pollution levels requires trajectory analysis, whereby both measured and modelled data is  
550 filtered to limit the data to that arriving from the clean sector, unaffected by land-based  
551 emission sources, as discussed in Section 2.2. Figure 10 displays trends in observed and  
552 simulated O<sub>3</sub> at Mace Head, separated into seasons and clean/EU influenced sectors according  
553 to the trajectory analysis. The figure shows that the clean sector has consistently higher O<sub>3</sub>  
554 concentrations than the EU influenced sector for Winter, Spring and Autumn, with the most  
555 significant disparity between clean and EU sectors in winter/spring when stratospheric  
556 intrusion and lightning NO<sub>x</sub> contribute most significantly to O<sub>3</sub>, as discussed in Section 3.4.2.  
557 O<sub>3</sub> originating from EU airmasses is susceptible to higher rates of dry deposition and removal  
558 via local pollution while traversing the land-mass westwards towards Mace Head, leading to  
559 higher O<sub>3</sub> in clean-sector air masses consistent with previous studies (Coleman et al., 2013). A  
560 decreasing trend in mean spring-time levels is observed for both clean and EU influenced  
561 sectors, consistent with the sustained decrease in precursor emissions in Europe and North  
562 America as displayed in Figure 9. An increasing trend (0.26 ppb/year) is observed in the winter-  
563 time EU influenced sector, with a trend of smaller magnitude (0.08 ppb/year) observed in the  
564 clean sector, indicating a decrease in winter-time O<sub>3</sub> depletion events due to decreasing

565 European emissions (from the EU influenced sector), consistent with the conclusions from  
566 previous studies of Mace Head surface O<sub>3</sub> (Derwent et al., 2024). Summer-time values do not  
567 exhibit a notable trend or a discrepancy between the clean and EU influenced sector  
568 measurements, indicating that there is little O<sub>3</sub> advected into Europe from the west in the  
569 summer months. It is noted that the seasonal trends exhibit slopes with p-values ranging  
570 between  $0.1 > p > 0.01$ , which denote trends of medium to high certainty, as defined by TOAR  
571 assessment criteria of Chang et al., 2023.

572 Although the model results display a negative bias for reasons outlined in section 3.4.1, the  
573 clean sector consistently exhibits higher simulated O<sub>3</sub> concentrations than the EU-influenced  
574 sector, except during the summer season. In the winter season, a significant increasing trend is  
575 observed for simulated O<sub>3</sub> from both sectors, aligning with observations. A decreasing trend is  
576 observed during the summer season, again consistent with the observations for both sectors. In  
577 spring and autumn, modelled O<sub>3</sub> exhibits a positive trend over the simulation period.

578 Trends in contributors of model O<sub>3</sub> during the different seasons for the clean and EU-influenced  
579 sector are shown in Tables S3 to S4 in the supplementary material. Aviation and East Asian  
580 NO<sub>x</sub> show consistently positive and significant trends in both sectors, while North America  
581 NO<sub>x</sub> shows strong negative trends throughout the year in the clean sector. For the EU-  
582 influenced sector in NO<sub>x</sub> tagged (Table S2), similar positive trends observed for aviation and  
583 East Asian, with North America NO<sub>x</sub> remaining negative and European NO<sub>x</sub> showing more  
584 significant declines in spring and winter. In case of VOC tagged O<sub>3</sub>, the East Asian VOCs  
585 shows an increasing trend and North America VOCs show a negative across all seasons in both  
586 sectors (Table S3 and S4). European VOCs also show a consistent negative trend, particularly  
587 strong in the EU-influenced sector. Methane trends are seasonally positive, especially in spring  
588 and winter.

589 It is observed that the model consistently simulates O<sub>3</sub> at lower concentrations than that  
 590 observed at Mace Head. This is not surprising, considering the coarse resolution of the model,  
 591 which limits its ability to represent fine-scale processes and dry deposition accurately. Dry  
 592 deposition is typically higher over land, and the grid cell covering Mace Head includes land  
 593 area, as shown in Figure S5 of the supplementary material. Further, as explained by Fiore et  
 594 al. (2009), models average the landscape characteristics within a grid cell, which can enhance  
 595 O<sub>3</sub> deposition and result in lower simulated O<sub>3</sub> concentrations; hence, the discrepancy is more  
 596 pronounced in the clean sector data.

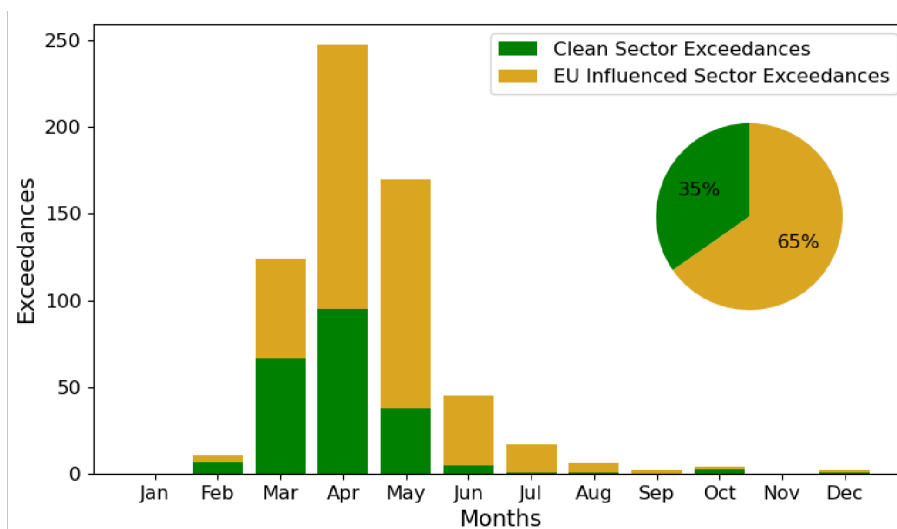


597  
 598 **Figure 10.** - Trend in seasonal Average of observed O<sub>3</sub> (black ) and Model O<sub>3</sub> (red) at Mace  
 599 head, separated into clean sector and EU-influenced sector.

600 **3.6 Exceedances from the clean and EU-influenced sector at Mace Head**

601 Exceedances observed at Mace Head between 2000 and 2022 are separated into clean and EU-  
602 influenced sectors based on trajectory air masses and shown in Figure 11. 35% of all  
603 exceedances for this period occurred in clean air masses, the remainder occurring when air  
604 masses include EU outflow and contribution from local sources to enhance surface O<sub>3</sub> at Mace  
605 Head, which is already elevated compared to inland and urban sites. It is notable that there is  
606 a higher proportion of exceedances that occur from EU- influenced sector, despite higher mean  
607 observations from the clean sector for all seasons. This occurs because of an enhancement of  
608 surface O<sub>3</sub> occur during an influx of polluted air from the EU, UK or local sources. The EU  
609 influence on exceedance becomes more proportionally prominent in late Spring and summer,  
610 with more frequent easterly airflow when there is a higher occurrence of stagnation events.

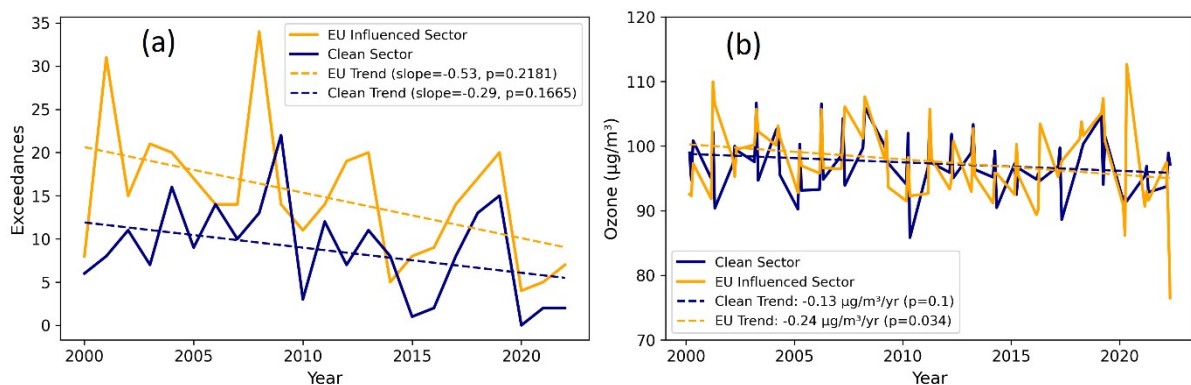
611



612

613 **Figure 11.** - Exceedances measured at Mace Head per month from 2000 until 2022, during  
614 the clean air sector(green) and EU influenced sector (yellow). The percentage of both to total  
615 exceedances is shown in the inlay.

616 Figure 12 shows the trend in spring-time exceedances and the 95<sup>th</sup> percentile Springtime O<sub>3</sub>  
617 measured at Mace Head between 2000 and 2022. A decreasing trend in exceedances is  
618 observed, with a greater decreasing trend from the EU and the locally influenced sector. This  
619 indicates that the changes that are driving the reduction in the exceedances in Europe are  
620 coming into effect at a quicker rate than the changes that are driving the O<sub>3</sub> event reduction  
621 over the North Atlantic. The trends in the exceedance counts are not significant, according  
622 to the criteria in Chang et al., 2023, but there is a statistically significant decreasing trend in  
623 the 95<sup>th</sup> percentile springtime surface O<sub>3</sub> over the measurement record. Figure 12 (b) shows  
624 the trend in spring-time surface O<sub>3</sub> measured at Mace Head segregated into the Clean and  
625 EU-influenced sector. The trend is more significant both in magnitude and statistical  
626 certainty for the EU-influenced sector, indicating that European emission changes have a  
627 more pronounced effect on springtime O<sub>3</sub> measured at Mace Head O<sub>3</sub> as compared to changes  
628 affecting O<sub>3</sub> transported or formed over the North Atlantic.



629

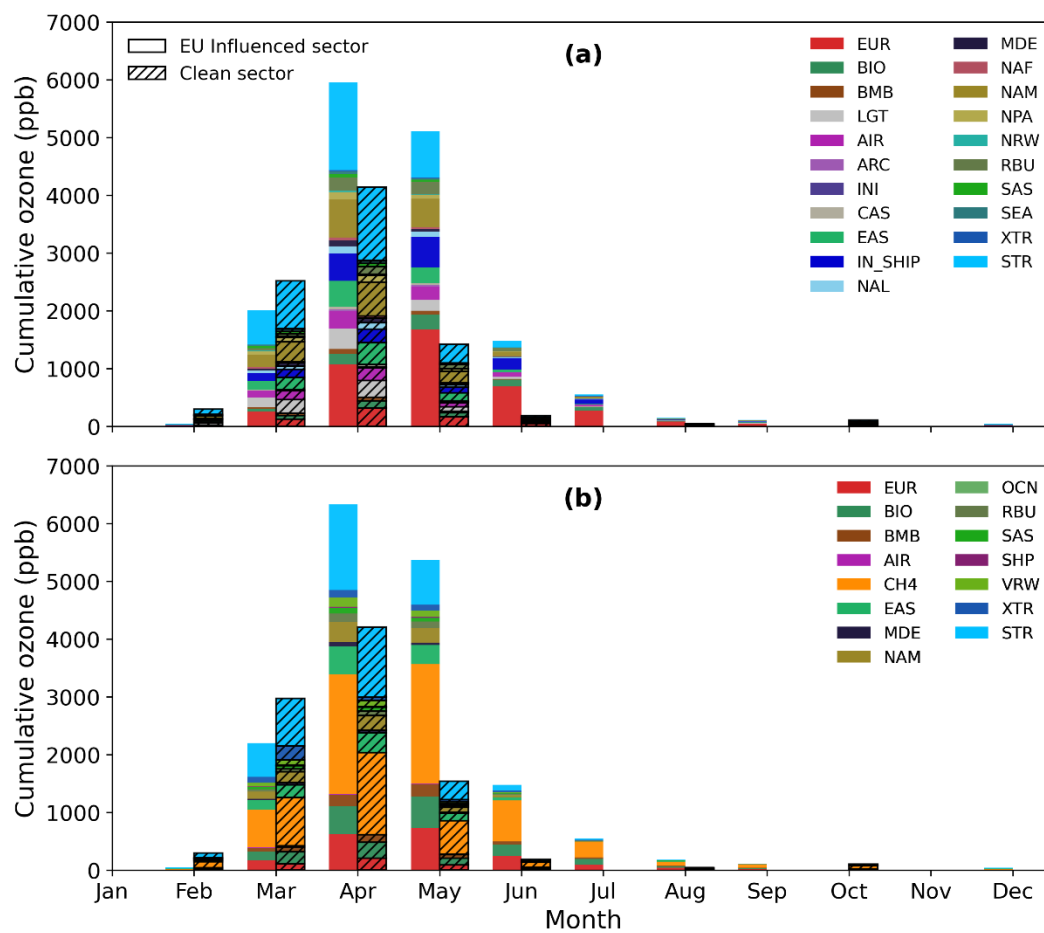
630 **Figure 12.** (a) The trend in Spring-time exceedances measured at Mace Head between 2000  
631 and 2022 (blue) with the clean-air exceedances (gold), and (b) The trend in the 95<sup>th</sup> percentile  
632 of spring (Mar- May) O<sub>3</sub> measured in µg /m<sup>3</sup> for the clean sector (blue) and the EU-influenced  
633 sector (gold).

634 Figure 13 shows monthly cumulative contributions to simulated O<sub>3</sub> concentrations within the  
635 Mace Head grid cell for NO<sub>x</sub> and VOC tagging during O<sub>3</sub> exceedance. First, the exceedances

636 were identified from O<sub>3</sub> observations as discussed in Section 3.3 and then these exceedances  
637 were divided into two sectors the EU-influenced sector and the clean sector. In Figure 13, the  
638 hourly O<sub>3</sub> exceedance cumulative concentrations in ppb, along with the contributions from  
639 different parameters, are shown. It indicates which parameters contribute more to the  
640 exceedances, in both the EU-influenced sector or the clean sector. In this, the µg/m<sup>3</sup>-to-ppb  
641 conversion is applied to the exceedances.

642 These exceedances are categorised into clean and EU-influenced sectors. The maximum  
643 exceedances are observed in March to May. From Figure 13 (a), it is clear that stratospheric  
644 intrusion, North American NO<sub>x</sub>, European NO<sub>x</sub>, and East Asian NO<sub>x</sub> are the major  
645 contributors driving exceedances at Mace Head during the spring months (March to May).  
646 European emissions dominate the supply of NO<sub>x</sub> precursors in April, reaching their peak in  
647 May. Figure 13 (b) shows that CH<sub>4</sub> is the most dominant VOC source, followed by stratospheric  
648 intrusion and Biomass burning. North American and European VOC emissions also contribute  
649 significantly to O<sub>3</sub> formation during this period. Collectively, these findings highlight the  
650 complex interplay of regional and global sources in driving surface O<sub>3</sub> exceedances over the  
651 Irish domain.

652 North American NO<sub>x</sub> also contributes significantly to exceedance in both clean and EU-  
653 influenced sectors at Mace Head during March to May month, likely due to long-range  
654 transport and mixing, regional stagnation or synoptic-scale recirculation. In the case of VOC  
655 tagging, stratospheric intrusion, and CH<sub>4</sub> show notable contributions. Biomass burning, East  
656 Asian emissions and North American VOC emissions also play a role in O<sub>3</sub> exceedances.  
657 The IN\_SHIP indicates the total contribution of all international shipping emissions  
658 (ENA,NAE,NAW,NPA,BNS,HBY,IDO,MBC,SHO) which is significant in April and May.



659  
 660 **Figure 13** -Monthly cumulative Mace Head grid cell O<sub>3</sub> contributions to EU influenced sector  
 661 and clean sector exceedances (a) NO<sub>x</sub> tagging and (b) VOC tagging Mace Head grid cell.

662 **4. Conclusion**

663 This study highlights the complexities of O<sub>3</sub> pollution in Ireland, revealing that coastal areas  
 664 experience higher O<sub>3</sub> concentrations than rural and urban environments, attributed to the effect  
 665 of transboundary pollution and stratospheric intrusion. Over the last two decades, urban sites  
 666 have shown a significant increasing trend in O<sub>3</sub> levels, particularly in winter, influenced by  
 667 decreasing anthropogenic pollution in Europe, the UK, North America and local to the  
 668 observation sites, representing a decline in chemical removal mechanism (Derwent et al., 2024;  
 669 Simmonds and Derwent, 1991). The analysis also points out that the majority of exceedances  
 670 at coastal monitoring sites coincide with the annual spring maxima. Using the advanced  
 671 capabilities of the CAM4-Chem model with dual NO<sub>x</sub> and VOC tagging, we identified key

672 factors affecting seasonal O<sub>3</sub> variations, such as the spring-time peak and summer dip, driven  
673 by a mix of stratospheric intrusion, NO<sub>x</sub> emissions from lightning, long range and hemispheric  
674 transport, regional emissions and aviation related emissions. Trend analysis from simulation  
675 results identified East Asian and aviation emissions as significant contributors to the rising  
676 winter trends in O<sub>3</sub>, while reductions in North American and European emissions accounted  
677 for the decrease in spring peaks. This study provides a comprehensive understanding of the  
678 various factors affecting O<sub>3</sub> levels in Ireland, offering important insights for the development  
679 of O<sub>3</sub> pollution control policies.

680

#### 681 **Data availability**

682 All data are available upon request.

683

#### 684 **Author contributions**

685 LC designed the study. NK analyzed the data and wrote the manuscript. TA and TB provided  
686 CAM-Chem model results and reviewed the manuscript. EC contributed to the air mass  
687 trajectories analysis. JO, CD and DM reviewed the manuscript and edited it. LC edited it  
688 with contributions from all coauthors .

#### 689 **Competing interests**

690 The authors declare that at least one of the authors sits on the editorial board of ACP.

691 **Acknowledgement** - The authors acknowledge the Environmental Protection Agency (EPA)  
692 of Ireland for their financial support of the Ozone project under the EPA Research  
693 Programme.

694 2021-2030 (project number 2022-CE-1133), and the European Union's Horizon Europe  
695 Research and Innovation programme under HORIZON-CL5-2022-D1-02 (grant no.

696 101081430-PARIS).

697

698 **References**

699 Aghedo, A. M., Bowman, K. W., Worden, H. M., Kulawik, S. S., Shindell, D. T., Lamarque,  
700 J. F., Faluvegi, G., Parrington, M., Jones, D. B. A., and Rast, S.: The vertical distribution of  
701 ozone instantaneous radiative forcing from satellite and chemistry climate models, *Journal*  
702 *of Geophysical Research Atmospheres*, 116, <https://doi.org/10.1029/2010JD014243>, 2011.

703 Anenberg, S. C., West, J. J., Fiore, A. M., Jaffe, D. A., Prather, M. J., Bergmann, D., Cuvelier,  
704 K., Dentener, F. J., Duncan, B. N., Gauss, M., Hess, P., Jonson, J. E., Lupu, A., Mackenzie,  
705 I. A., Marmer, E., Park, R. J., Sanderson, M. G., Schultz, M., Shindell, D. T., Szopa, S.,  
706 Vivanco, M. G., Wild, O., and Zeng, G.: Intercontinental impacts of ozone pollution on  
707 human mortality, *Environ Sci Technol*, 43, 6482–6487, <https://doi.org/10.1021/es900518z>,  
708 2009.

709 Ansari, T., Nalam, A., Lupaşcu, A., Hinz, C., Grasse, S., & Butler, T.: Explaining trends and  
710 changing seasonal cycles of surface ozone in North America and Europe over the 2000–2018  
711 period: a global modelling study with NO<sub>x</sub> and VOC tagging, *Atmospheric Chemistry and*  
712 *Physics*, 25, 16833–16833. <https://doi.org/10.5194/acp-25-16833-2025>.

713 Archibald, A. T., Neu, J. L., Elshorbany, Y. F., Cooper, O. R., Young, P. J., Frost, G. J.,  
714 Galbally, I. E., Gerosa, G., Granier, C., and Griffiths, P. T.: Tropospheric Ozone Assessment  
715 Report : A critical review of changes in the tropospheric ozone burden and budget from 1850  
716 to 2100, 1– 53, 2020.

717 Ashmore, M. R.: Assessing the future global impacts of ozone on vegetation, *Plant Cell*  
718 *Environ*, 28, 949–964, <https://doi.org/10.1111/j.1365-3040.2005.01341.x>, 2005.

719 Atkinson-Palombo, C. M., Miller, J. A., and Balling, R. C.: Quantifying the ozone “weekend  
720 effect” at various locations in Phoenix, Arizona, *Atmos Environ*, 40, 7644–7658,  
721 <https://doi.org/10.1016/j.atmosenv.2006.05.023>, 2006.

722 Auvray, M. and Bey, I.: Long-range transport to Europe: Seasonal variations and implications  
723 for the European ozone budget, <https://doi.org/10.1029/2004JD005503>, 16 June 2005.

724 Bessagnet, B., Pirovano, G., Mircea, M., Cuvelier, C., Aulinger, A., Calori, G., Ciarelli, G.,  
725 Manders, A., Stern, R., Tsyro, S., Garcíá Vivanco, M., Thunis, P., Pay, M. T., Colette, A.,  
726 Couvidat, F., Meleux, F., Rouïl, L., Ung, A., Aksoyoglu, S., Baldasano, J. M., Bieser, J.,  
727 Briganti, G., Cappelletti, A., D’Isidoro, M., Finardi, S., Kranenburg, R., Silibello, C.,  
728 Carnevale, C., Aas, W., Dupont, J. C., Fagerli, H., Gonzalez, L., Menut, L., Prévôt, A. S. H.,  
729 Roberts, P., and White, L.: Presentation of the EURODELTA III intercomparison  
730 exercise evaluation of the chemistry transport models’ performance on criteria pollutants and  
731 joint analysis with meteorology, *Atmos Chem Phys*, 16, 12667–12701,  
732 <https://doi.org/10.5194/acp-16-12667-2016>, 2016.

733 Bonaccorso, B., Peres, D. J., Cancelliere, A., and Di Mauro, G.: Probabilistic forecasting of  
734 drought class transitions in Sicily (Italy) using Standardized Precipitation Index and North  
735 Atlantic Oscillation Index, *Hydrol. Earth Syst. Sci.*, 19, 389–404, 2015.

736 Boylan, P., Helmig, D., and Oltmans, S.: Ozone in the Atlantic Ocean marine boundary layer,  
737 1–13, <https://doi.org/10.12952/journal.elementa.000045>, 2014.

738 Butchart, N., Charlton-Perez, A. J., Cionni, I., Hardiman, S. C., Haynes, P. H., Krüger, K.,  
739 Kushner, P. J., Newman, P. A., Osprey, S. M., Perlwitz, J., Sigmund, M., Wang, L., Akiyoshi,  
740 H., Austin, J., Bekki, S., Baumgaertner, A., Braesicke, P., Brhl, C., Chipperfield, M.,  
741 Dameris, M., Dhomse, S., Eyring, V., Garcia, R., Garny, H., Jöckel, P., Lamarque, J. F.,

742 Marchand, M., Michou, M., Morgenstern, O., Nakamura, T., Pawson, S., Plummer, D., Pyle,  
743 J., Rozanov, E., Scinocca, J., Shepherd, T. G., Shibata, K., Smale, D., Teyssèdre, H., Tian,  
744 W., Waugh, D., and Yamashita, Y.: Multimodel climate and variability of the stratosphere,  
745 *Journal of Geophysical Research Atmospheres*, 116, <https://doi.org/10.1029/2010JD014995>,  
746 2011.

747 Butler, T., Lupascu, A., Coates, J., and Zhu, S.: TOAST 1.0: Tropospheric ozone attribution  
748 of sources with tagging for CESM 1.2.2, *Geosci Model Dev*,  
749 11, 2825–2840, <https://doi.org/10.5194/gmd-11-2825-2018>, 2018.

750 Butler, T., Lupascu, A., and Nalam, A.: Attribution of ground-level ozone to anthropogenic  
751 and natural sources of nitrogen oxides and reactive carbon in a global chemical transport  
752 model, *Atmos Chem Phys*, 20, 10707–10731, <https://doi.org/10.5194/acp-20-10707-2020>,  
753 2020.

754 Carslaw, D. C.: On the changing seasonal cycles and trends of ozone at Mace Head, Ireland,  
755 *Atmos. Chem. Phys*, 3441–3450 pp., 2005.

756 Chang, K.-L., Schultz, M. G., Koren, G., & Selke, N.: Guidance note on best statistical 1108  
757 practices for TOAR analyses. arXiv preprint arXiv:2304.14236, 2023.

758 Coates, J., Mar, K. A., Ojha, N., & Butler, T. M.: The influence of temperature on ozone  
759 production under varying NO<sub>x</sub> conditions—a modelling study. *Atmospheric Chemistry and*  
760 *Physics*, 16(18), 11601-11615, 2016.

761 Coleman, L., McVeigh, P., Berresheim, H., Martino, M., and O’Dowd, C. D.: Photochemical  
762 impact on ozone fluxes in coastal waters, *Advances in Meteorology*, 2012,  
763 <https://doi.org/10.1155/2012/943785>, 2012.

764 Coleman, L., Martin, D., Varghese, S., Jennings, S. G., and O'Dowd, C. D.: Assessment of  
765 changing meteorology and emissions on air quality using a regional climate model: Impact  
766 on ozone, *Atmos Environ*, 69, 198–210, <https://doi.org/10.1016/j.atmosenv.2012.11.048>,  
767 2013.

768 Creilson, J. K., Fishman, J., & Wozniak, A. E. Intercontinental transport of tropospheric  
769 ozone: A study of its seasonal variability across the North Atlantic utilizing tropospheric  
770 ozone residuals and its relationship to the North Atlantic Oscillation. *Atmospheric Chemistry  
771 and Physics*, 3(6), 2053–2066. <https://doi.org/10.5194/acp-3-2053-2003>,2003.

772 Crippa, M., Guizzardi, D., Butler, T. M., Keating, T., Wu, R., Kaminski, J., Kiesewetter, G.,  
773 Im, U., Bessagnet, B., Chabrillat, S., et al.: The HTAP\_v3 emission mosaic: merging regional  
774 and global monthly emissions (2000–2018) to support air quality modelling and policies, *Earth  
775 Syst. Sci. Data*, 15, 2667–2696, 2023.

776 Crippa, M., Guizzardi, D., Pagani, F., Schiavina, M., Melchiorri, M., Pisoni, E., Graziosi, F.,  
777 Muntean, M., Maes, J., Dijkstra, L., Van Damme, M., Clarisse, L., and Coheur, P.: Insights  
778 into the spatial distribution of global, national, and subnational greenhouse gas emissions in  
779 the Emissions Database for Global Atmospheric Research (EDGAR v8.0), *Earth Syst Sci  
780 Data*, 16, 2811–2830, <https://doi.org/10.5194/essd-16-2811-2024>, 2024.

781 Derwent, R. G.: Observation and interpretation of the seasonal cycles in the surface  
782 concentrations of ozone and carbon monoxide at Mace Head, Ireland from 1990 to 1994,  
783 *Atmos Environ*, 32, 2310, [https://doi.org/10.1016/S1352-2310\(97\)00338-5](https://doi.org/10.1016/S1352-2310(97)00338-5), 1998.

784 Derwent, R. G., Simmonds, P. G., and Collins, W. J.: Ozone and carbon monoxide  
785 measurements at a remote maritime location, mace head, Ireland, from 1990 to 1992, *Atmos  
786 Environ*, 28, 2623–2637, [https://doi.org/10.1016/1352-2310\(94\)90436-7](https://doi.org/10.1016/1352-2310(94)90436-7), 1994.

787 Derwent, R. G., Collins, W. J., Johnson, C. E., and Stevenson, D. S.: Transient behaviour of  
788 tropospheric ozone precursors in a global 3-d ctm and their indirect greenhouse effects, 2001.

789 Derwent, R. G., Stevenson, D. S., Collins, W. J., and Johnson, C. E.: Intercontinental  
790 transport and the origins of the ozone observed at surface sites in Europe, *Atmos Environ*,  
791 38, 1891–1901, <https://doi.org/10.1016/j.atmosenv.2004.01.008>, 2004.

792 Derwent, R. G., Stevenson, D. S., Doherty, R. M., Collins, W. J., and Sanderson, M. G.: How  
793 is surface ozone in Europe linked to Asian and North American NO<sub>x</sub> emissions?, *Atmos*  
794 *Environ*, 42, 7412–7422, <https://doi.org/10.1016/j.atmosenv.2008.06.037>, 2008.

795 Derwent, R. G., Manning, A. J., Simmonds, P. G., Spain, T. G., and O’Doherty, S.: Analysis  
796 and interpretation of 25 years of ozone observations at the Mace Head Atmospheric Research  
797 Station on the Atlantic Ocean coast of Ireland from 1987 to 2012, *Atmos Environ*, 80, 361–  
798 368, <https://doi.org/10.1016/j.atmosenv.2013.08.003>, 2013.

799 Derwent, R. G., Manning, A. J., Simmonds, P. G., and Doherty, S. O.: Long-term trends in  
800 ozone in baseline and European regionally-polluted air at Mace Head, Ireland over a 30-year  
801 period, *Atmos Environ*, 179, 279–287, <https://doi.org/10.1016/j.atmosenv.2018.02.024>,  
802 2018.

803 Derwent, R. G., Parrish, D. D., Manning, A. J., Spain, T. G., Simmonds, P. G., and O’Doherty,  
804 S.: Ozone at Mace Head, Ireland from 1987 to 2021: Declining baselines, phase-out of  
805 European regional pollution, COVID-19 impacts, *Atmos Environ*, 320,  
806 <https://doi.org/10.1016/j.atmosenv.2023.120322>, 2024.

807 Ding, J., Dai, Q., Fan, W., Lu, M., Zhang, Y., Han, S., and Feng, Y.: Impacts of meteorology  
808 and precursor emission change on O<sub>3</sub> variation in Tianjin, China from 2015 to 2021, *J*  
809 *Environ Sci (China)*, 126, 506–516, <https://doi.org/10.1016/j.jes.2022.03.010>, 2023.

810 Donlon, B., Cahalane, A., and Fanning, A.: Ireland's State of the Environment Report 2024  
811 Editors, 2024.

812 Draxler, R. R. Evaluation of an Ensemble Dispersion Calculation.  
813 <http://wesley.wwb.noaa.gov/reanalysis.html>,2003.

814 EEA: Exceedance of air quality standards in Europe (Indicator AIR003, Issue.1109)2024  
815 <https://www.eea.europa.eu/en/analysis/indicators/exceedance-of-air-quality-standards>

816 Emmons, L. K., Hess, P. G., Lamarque, J.-F., and Pfister, G. G.: Tagged ozone mechanism for  
817 MOZART-4, CAM-chem and other chemical transport models, *Atmos. Chem. Phys.*, **12**, 4607–  
818 4622, <https://doi.org/10.5194/acp-12-4607-2012>,2012.

819 Finch, D. P., & Palmer, P. I., Increasing ambient surface ozone levels over the UK accompanied  
820 by fewer extreme events. *Atmospheric Environment*, **237**, 117627,  
821 <https://doi.org/https://doi.org/10.1016/j.atmosenv.2020.117627>, 2020.

822 Fiore, A. M., Dentener, F. J., Wild, O., Cuvelier, C., Schultz, M. G., Hess, P., Textor, C.,  
823 Schulz, M., Doherty, R. M., Horowitz, L. W., MacKenzie, I. A., Sanderson, M. G., Shindell,  
824 D. T., Stevenson, D. S., Szopa, S., Van Dingenen, R., Zeng, G., Atherton, C., Bergmann, D.,  
825 Bey, I., Carmichael, G., Collins, W. J., Duncan, B. N., Faluvegi, G., Folberth, G., Gauss, M.,  
826 Gong, S., Hauglustaine, D., Holloway, T., Isaksen, I. S. A., Jacob, D. J., Jonson, J. E.,  
827 Kaminski, J. W., Keating, T. J., Lupu, A., Manner, E., Montanaro, V., Park, R. J., Pitari, G.,  
828 Pringle, K. J., Pyle, J. A., Schroeder, S., Vivanco, M. G., Wind, P., Wojcik, G., Wu, S., and  
829 Zuber, A.: Multimodel estimates of intercontinental source-receptor relationships for ozone  
830 pollution, *Journal of Geophysical Research Atmospheres*, **114**,  
831 <https://doi.org/10.1029/2008JD010816>, 2009.

832 Fowler, D., Coyle, M., Skiba, U., Sutton, M. A., Cape, J. N., Reis, S., Sheppard, L. J., Jenkins,  
833 A., Grizzetti, B., Galloway, J. N., Vitousek, P., Leach, A., Bouwman, A. F., Butterbach-Bahl,  
834 K., Dentener, F., Stevenson, D., Amann, M., and Voss, M.: The global nitrogen cycle in  
835 theTwentyfirst century, *Philosophical Transactions of the Royal Society B: Biological*  
836 *Sciences*, 368, <https://doi.org/10.1098/rstb.2013.0164>, 2013.

837 Girach, I. A., Tripathi, N., Nair, P. R., Sahu, L. K., and Ojha, N.: O<sub>3</sub> and CO in the South  
838 Asian outflow over the Bay of Bengal: Impact of monsoonal dynamics and chemistry, *Atmos*  
839 *Environ*, 233, <https://doi.org/10.1016/j.atmosenv.2020.117610>, 2020.

840 Griffiths, P. T., Murray, L. T., Zeng, G., Shin, Y. M., Abraham, N. L., Archibald, A. T., Deushi,  
841 M., Emmons, L. K., Galbally, I. E., Hassler, B., Horowitz, L. W., Keeble, J., Liu, J., Moeini,  
842 O., Naik, V., and Connor, F. M. O.: Tropospheric ozone in CMIP6 simulations, 4187–4218,  
843 2021.

844 Grigas, T., Ovadnevaite, J., Ceburnis, D., Moran, E., McGovern, F. M., Jennings, S. G., and  
845 O’Dowd, C.: Sophisticated Clean Air Strategies Required to Mitigate Against Particulate  
846 Organic Pollution, *Sci Rep*, 7, <https://doi.org/10.1038/srep44737>, 2017.

847 Guerreiro, C. B. B., Foltescu, V., and de Leeuw, F.: Air quality status and trends in Europe,  
848 *Atmos Environ*, 98, 376–384, <https://doi.org/10.1016/j.atmosenv.2014.09.017>, 2014.

849 Huang, M., Carmichael, G. R., Pierce, R. B., Jo, D. S., Park, R. J., Flemming, J., Emmons, L.  
850 K., Bowman, K. W., Henze, D. K., Davila, Y., Sudo, K., Jonson, J. E., Lund, M. T., Janssens-  
851 Maenhout, G., Dentener, F. J., Keating, T. J., & Saiz-Lopez, A. (2017). Impact of  
852 intercontinental pollution transport on North American ozone air pollution: An HTAP phase 2  
853 multi-model study. *Atmospheric Chemistry and Physics*, 17(9), 5721–5750.  
854 <https://doi.org/10.5194/acp-17-5721-2017>

855 IPCC, Masson-Delmotte, V., Zhai, P., Chen, Y., Goldfarb, L., Gomis, M. I., Matthews, J. B.  
856 R., Berger, S., Huang, M., Yelekçi, O., Yu, R., Zhou, B., Lonnoy, E., Maycock, T. K.,  
857 Waterfield, T., Leitzell, K., & Caud, N. Working Group I Contribution to the Sixth  
858 Assessment Report of the Intergovernmental Panel on Climate Change Edited by.  
859 [www.ipcc.ch](http://www.ipcc.ch),2021.

860 Jeon, W. B., Lee, S. H., Lee, H., Park, C., Kim, D. H., and Park, S. Y.: A study on high ozone  
861 formation mechanism associated with change of NO<sub>x</sub>/VOCs ratio at a rural area in the  
862 Korean Peninsula, *Atmos Environ*, 89, 10–21,  
863 <https://doi.org/10.1016/j.atmosenv.2014.02.005>, 2014.

864 Jonson, J. E., Stohl, A., Fiore, A. M., Hess, P., Szopa, S., Wild, O., Zeng, G., Dentener, F. J.,  
865 Lupu, A., Schultz, M. G., Duncan, B. N., Sudo, K., Wind, P., Schulz, M., Marmer, E.,  
866 Cuvelier, C., Keating, T., Zuber, A., Valdebenito, A., Dorokhov, V., De Backer, H., Davies,  
867 J., Chen, G. H., Johnson, B., Tarasick, D. W., Stübi, R., Newchurch, M. J., Von Der Gathen,  
868 P., Steinbrecht, W., and Claude, H.: A multi-model analysis of vertical ozone profiles, *Atmos*  
869 *Chem Phys*, 10, 5759–5783, <https://doi.org/10.5194/acp-10-5759-2010>, 2010.

870

871 Khiem, M., Ooka, R., Huang, H., Hayami, H., Yoshikado, H., and Kawamoto, Y.: Analysis  
872 of the Relationship between Changes in Meteorological Conditions and the Variation in  
873 Summer Ozone Levels over the Central Kanto Area, *Advances in Meteorology*, 2010, 1–13,  
874 <https://doi.org/10.1155/2010/349248>, 2010.

875 Lamarque, J. F. and Solomon, S.: Impact of changes in climate and halocarbons on recent  
876 lower stratosphere ozone and temperature trends, *J Clim*, 23, 2599–  
877 2611, <https://doi.org/10.1175/2010JCLI3179.1>, 2010.

878 Lamarque, J. F., Kinnison, D. E., Hess, P. G., and Vitt, F. M.: Simulated lower stratospheric  
879 trends between 1970 and 2005: Identifying the role of climate and composition changes,  
880 *Journal of Geophysical Research Atmospheres*, 113, <https://doi.org/10.1029/2007JD009277>,  
881 2008.

882 Lamarque, J. F., Bond, T. C., Eyring, V., Granier, C., Heil, A., Klimont, Z., Lee, D., Liousse,  
883 C., Mieville, A., Owen, B., Schultz, M. G., Shindell, D., Smith, S. J., Stehfest, E., Van  
884 Aardenne, J., Cooper, O. R., Kainuma, M., Mahowald, N., McConnell, J. R., Naik, V., Riahi,  
885 K., and Van Vuuren, D. P.: Historical (1850-2000) gridded anthropogenic and biomass  
886 burning emissions of reactive gases and aerosols: Methodology and application, *Atmos*  
887 *Chem Phys*, 10, 7017–7039, <https://doi.org/10.5194/acp-10-7017-2010>, 2010.

888 Lamarque, J. F., Emmons, L. K., Hess, P. G., Kinnison, D. E., Tilmes, S., Vitt, F., Heald, C.  
889 L., Holland, E. A., Lauritzen, P. H., Neu, J., Orlando, J. J., Rasch, P. J., and Tyndall, G. K.:  
890 CAMchem: Description and evaluation of interactive atmospheric chemistry in the  
891 Community Earth System Model, *Geosci Model Dev*, 5, 369–411,  
892 <https://doi.org/10.5194/gmd-5-369-2012>, 2012.

893

894 Lefohn, A. S., Malley, C. S., Smith, L., Wells, B., Hazucha, M., Simon, H., Naik, V., Mills,  
895 G., Schultz, M. G., Paoletti, E., De Marco, A., Xu, X., Zhang, L., Wang, T., Neufeld, H. S.,  
896 Musselman, R. C., Tarasick, D., Brauer, M., Feng, Z., Tang, H., Kobayashi, K., Sicard, P.,  
897 Solberg, S., and Gerosa, G.: Tropospheric ozone assessment report: Global ozone metrics for  
898 climate change, human health, and crop/ecosystem research,  
899 *Elementa*, 6, <https://doi.org/10.1525/elementa.279>, 2018.

900 Lin, M., Fiore, A. M., Cooper, O. R., Horowitz, L. W., Langford, A. O., Levy, H., Johnson,  
901 B. J., Naik, V., Oltmans, S. J., and Senff, C. J.: Springtime high surface ozone events over

902 the western United States: Quantifying the role of stratospheric intrusions, *Journal of*  
903 *Geophysical Research Atmospheres*, 117, <https://doi.org/10.1029/2012JD018151>, 2012.

904 Lin, Y., Jiang, F., Zhao, J., Zhu, G., He, X., Ma, X., Li, S., Sabel, C. E., and Wang, H.: Impacts  
905 of O<sub>3</sub> on premature mortality and crop yield loss across China, *Atmos. Environ.*, **194**, 41–47,  
906 <https://doi.org/10.1016/j.atmosenv.2018.09.024>, 2018.

907 Lupaşcu, A., Otero, N., Minkos, A., and Butler, T.: Attribution of surface ozone to NO<sub>x</sub> and  
908 volatile organic compound sources during two different high ozone events, *Atmos Chem*  
909 *Phys*, 22, 11675–11699, <https://doi.org/10.5194/acp-22-11675-2022>, 2022.

910 McVeigh, P., O'Dowd, C., and Berresheim, H.: Eddy Correlation Measurements of Ozone  
911 Fluxes over Coastal Waters West of Ireland, *Advances in Meteorology*, 2010,  
912 <https://doi.org/10.1155/2010/754941>, 2010.

913 Moiseenko, K. B., Vasileva, A. V., Skorokhod, A. I., Belikov, I. B., Repin, A. Y., and  
914 Shtabkin, Y. A.: Regional Impact of Ozone Precursor Emissions on NO<sub>x</sub> and O<sub>3</sub> Levels at  
915 ZOTTO Tall Tower in Central Siberia, *Earth and Space Science*, 8,  
916 <https://doi.org/10.1029/2021EA001762>, 2021.

917 Molod, A., Takacs, L., Suarez, M., and Bacmeister, J.: Development of the GEOS-5  
918 atmospheric general circulation model: Evolution from MERRA to MERRA2, *Geosci Model*  
919 *Dev*, 8, 1339–1356, <https://doi.org/10.5194/gmd-8-1339-2015>, 2015.

920 Monks, P. S., Archibald, A. T., Colette, A., Cooper, O., Coyle, M., Derwent, R., Fowler, D.,  
921 Granier, C., Law, K. S., Mills, G. E., Stevenson, D. S., Tarasova, O., Thouret, V., Von  
922 Schneidemesser, E., Sommariva, R., Wild, O., and Williams, M. L.: Tropospheric ozone and  
923 its precursors from the urban to the global scale from air quality to short-lived climate forcer,  
924 <https://doi.org/10.5194/acp-15-8889-2015>, 13 August 2015.

925 Nalam, A., Lupaşcu, A., Ansari, T., and Butler, T.: Regional and sectoral contributions of NO<sub>x</sub>  
926 and reactive carbon emission sources to global trends in tropospheric ozone during the 2000–  
927 2018 period, *Atmos. Chem. Phys.*, 25, 5287–5311, <https://doi.org/10.5194/acp-25-5287-2025>,  
928 2025.

929 Nelson, B. S., et al.: Urban ozone trends in Europe and the USA (2000–2021), *Atmos. Chem.*  
930 *Phys.*, 25, 16009–16032, <https://doi.org/10.5194/acp-25-16009-2025>, 2025

931 O’Dowd, C., Ceburnis, D., Ovadnevaite, J., Vaishya, A., Rinaldi, M., and Facchini, M. C.:  
932 Do anthropogenic, continental or coastal aerosol sources impact on a marine aerosol  
933 signature at Mace Head?, *Atmos Chem Phys*, 14, 10687–10704, <https://doi.org/10.5194/acp->  
934 [14-10687-2014](https://doi.org/10.5194/acp-14-10687-2014), 2014.

935 Oltmans, S.J., Lefohn, A.S., Harris, J.M., Galbally, I., Scheel, H.E., Bodeker, G., Brunke, E.,  
936 Claude, H., Tarasick, D., Johnson, B.J. and Simmonds, P.: Long-term changes in tropospheric  
937 ozone. *Atmos Environ*, 40(17), 3156-3173, <https://doi.org/10.1016/j.atmosenv.2006.01.029>. ,  
938 2006.

939 Oltmans, S. J., Lefohn, A. S., Shadwick, D., Harris, J. M., Scheel, H. E., Galbally, I., Tarasick,  
940 D. W., Johnson, B. J., Brunke, E. G., Claude, H., Zeng, G., Nichol, S., Schmidlin, F., Davies,  
941 J., Cuevas, E., Redondas, A., Naoe, H., Nakano, T., and Kawasato, T.: Recent tropospheric  
942 ozone changes - A pattern dominated by slow or no growth, *Atmos Environ*, 67, 331–351,  
943 <https://doi.org/10.1016/j.atmosenv.2012.10.057>, 2013.

944 Ordóñez, C., Garrido-Perez, J. M., & García-Herrera, R. Early spring near-surface ozone in  
945 Europe during the COVID-19 shutdown: Meteorological effects outweigh emission changes.  
946 *Science of The Total Environment*, 747, 141322.  
947 <https://doi.org/https://doi.org/10.1016/j.scitotenv.2020.141322>, 2020.

948 Pallé, E., and Butler, C.J.: "Comparison of sunshine records and synoptic cloud observations:  
949 a case study for Ireland." *Physics and Chemistry of the Earth, Parts A/B/C* 27.6-8: 405-414,  
950 2002.

951 Pan, C., Zhu, B., Gao, J., Hou, X., Kang, H., and Wang, D.: "Quantifying Arctic lower  
952 stratospheric ozone sources in winter and spring," *Sci Rep*, 8,  
953 <https://doi.org/10.1038/s41598018-27045-5>, 2018.

954 Paoletti, E.: "Impact of ozone on Mediterranean forests: A review," *Environmental Pollution*,  
955 144, 463–474, <https://doi.org/10.1016/j.envpol.2005.12.051>, 2006.

956 Pausata, F. S. R., Pozzoli, L., Vignati, E., & Dentener, F. J.: "North Atlantic Oscillation and  
957 tropospheric ozone variability in Europe: Model analysis and measurements  
958 intercomparison." *Atmospheric Chemistry and Physics*, 12(14), 6357–6376.  
959 <https://doi.org/10.5194/acp-126357-2012>, 2012.

960 Pio, C. A., Feliciano, M. S., Vermeulen, A. T., and Sousa, E. C.: "Seasonal variability of ozone  
961 dry deposition under southern European climate conditions, in Portugal," *Atmospheric  
962 Environment*, 195–205 pp., 2000.

963 Russo, M. R., Kerridge, B. J., Abraham, N. L., Keeble, J., Latter, B. G., Siddans, R., Weber,  
964 J., Griffiths, P. T., Pyle, J. A., and Archibald, A. T.: "Seasonal, interannual and decadal  
965 variability of tropospheric ozone in the North Atlantic: comparison of UM-UKCA and  
966 remote sensing observations for 2005 – 2018," 6169–6196, 2023.

967 Saiz-Lopez, A., Mahajan, A. S., Abbatt, J., Bobrowski, N., Brown, S. S., Burrows, J. P.,  
968 Carpenter, L. J., Chipperfield, M. P., Cuevas, C. A., Fernandez, R., Hossaini, R., Kinnison,  
969 D. E., Lamarque, J.-F., Finlayson-Pitts, B. J., Plane, J. M. C., Platt, U., Pratt, K., Ravishankara,  
970 A. R., Salawitch, R. J., Saltzman, E. S., Simpson, W. R., Solomon, S., Thornton, J. A., & Wang,

971 T.: The influence of short-lived halogens on atmospheric chemistry and climate, *Nature*, 648,  
972 289–299, <https://doi.org/10.1038/s41586-025-09753-x>, 2025.

973 Seinfeld, J. H. and Pandis, S. N.: *Atmospheric Chemistry and Physics: From Air Pollution to*  
974 *Climate Change*, 3rd edn., Wiley, Hoboken, NJ, 2016.

975 Shindell, D. T., Chin, M., Dentener, F., Doherty, R. M., Faluvegi, G., Fiore, A. M., Hess, P.,  
976 Koch, D. M., Mackenzie, I. A., Sanderson, M. G., Schultz, M. G., Schulz, M., Stevenson, D.  
977 S., Teich, H., Textor, C., Wild, O., Bergmann, D. J., Bey, I., Bian, H., Cuvelier, C., Duncan,  
978 B. N., Folberth, G., Horowitz, L. W., Jonson, J., Kaminski, J. W., Marmer, E., Park, R.,  
979 Pringle, K. J., Schroeder, S., Szopa, S., Takemura, T., Zeng, G., Keating, T. J., and Zuber, A.:  
980 *Atmospheric Chemistry and Physics A multi-model assessment of pollution transport to the*  
981 *Arctic*, *Atmos. Chem. Phys.*, 5353–5372 pp., 2008.

982 Sicard, P., Serra, R., and Rossello, P.: Spatiotemporal trends in ground-level ozone  
983 concentrations and metrics in France over the time period 1999-2012, *Environ Res.*, 149,  
984 122– 144, <https://doi.org/10.1016/j.envres.2016.05.014>, 2016.

985 Simmonds, P. G., Derwent, R. G., Manning, A. L., and Spain, G.: Significant growth in  
986 surface ozone at Mace Head, Ireland, 1987-2003, *Atmos Environ.*,  
987 38, 4769–4778, <https://doi.org/10.1016/j.atmosenv.2004.04.036>, 2004.

988 Soares, J., Plass, D., Kienzler, S., González Ortiz, A., Gsella, A., Horálek, J. (2023). *Health*  
989 *Risk Assessment of Air Pollution: assessing the environmental burden of disease in Europe*  
990 *in 2021 (Eionet Report – ETC HE 2023/7)*. European Topic Centre on Human Health and  
991 *the Environment*. <https://www.eionet.europa.eu/etcs/all-etc-reports>

992 Spohn, T. K., Martin, D., Geever, M., and O’Dowd, C.: Effect of COVID-19 lockdown on  
993 regional pollution in Ireland, *Air Qual Atmos Health*, 15,  
994 221–234, <https://doi.org/10.1007/s11869-021-01098-4>, 2022.

995 Stein, A. F., Draxler, R. R., Rolph, G. D., Stunder, B. J. B., Cohen, M. D., and Ngan, F.: NOAA’s  
996 HYSPLIT atmospheric transport and dispersion modeling system, *Bulletin of the American*  
997 *Meteorological Society*, 96, 2059–2077, <https://doi.org/10.1175/BAMS-D-14-00110.1>, 2015.

998 Stunder, B.J.B.: Global Data Assimilation System (GDAS) Archive Information, NOAA Air  
999 Resources Laboratory, Silver Spring, MD, USA, December 1, 2004

1000 Sudo, K., and H. Akimoto, H.: Global source attribution of tropospheric ozone: Long-range  
1001 transport from various source regions, *J. Geophys. Res.*, 112, D12302,  
1002 doi:[10.1029/2006JD007992](https://doi.org/10.1029/2006JD007992), 2007.

1003 Tan, J., Fu, J. S., Dentener, F., Sun, J., Emmons, L., Tilmes, S., Flemming, J., Takemura, T.,  
1004 Bian, H., Zhu, Q., Yang, C. E., and Keating, T.: Source contributions to sulfur and nitrogen  
1005 deposition - An HTAP II multi-model study on hemispheric transport, *Atmos Chem Phys*,  
1006 18, 12223–12240, <https://doi.org/10.5194/acp-18-12223-2018>, 2018.

1007 Tavella, R. A. and da Silva Júnior, F. M. R.: Watch out for trends: did ozone increased or  
1008 decreased during the COVID-19 pandemic?, *Environmental Science and Pollution Research*,  
1009 28, 67880–67885, <https://doi.org/10.1007/s11356-021-17142-w>, 2021.

1010 Tilmes, S., Sanderson, B. M., and O’Neill, B. C.: Climate impacts of geoengineering in a  
1011 delayed mitigation scenario, *Geophys Res Lett*, 43, 8222–  
1012 8229, <https://doi.org/10.1002/2016GL070122>, 2016.

1013 Tilmes, S., Lamarque, J. F., Emmons, L. K., Kinnison, D. E., Ma, P. L., Liu, X., Ghan, S.,  
1014 Bardeen, C., Arnold, S., Deeter, M., Vitt, F., Ryerson, T., Elkins, J. W., Moore, F., Spackman,

1015 J. R., and Val Martin, M.: Description and evaluation of tropospheric chemistry and aerosols  
1016 in the Community Earth System Model (CESM1.2), *Geosci Model Dev*, 8, 1395–1426,  
1017 <https://doi.org/10.5194/gmd-8-1395-2015>, 2015.

1018 Todorović, M. N., Radenković, M. B., Rajšić, S. F., & Ignjatović, L. M. Evaluation of  
1019 mortality attributed to air pollution in the three most populated cities in Serbia. *International*  
1020 *Journal of Environmental Science and Technology*, 16(11), 7059–7070.  
1021 <https://doi.org/10.1007/s13762019-02384-6>, 2019.

1022 Tripathi, O. P., Jennings, S. G., O’Dowd, C. D., Coleman, L., Leinert, S., O’Leary, B., Moran,  
1023 E., O’Doherty, S. J., and Spain, T. G.: Statistical analysis of eight surface ozone measurement  
1024 series for various sites in Ireland, *Journal of Geophysical Research Atmospheres*, 115, 1–20,  
1025 <https://doi.org/10.1029/2010JD014040>, 2010.

1026 Tripathi, O. P., Jennings, S. G., O’Dowd, C., O’Leary, B., Lambkin, K., Moran, E.,  
1027 O’Doherty, S. J., and Gerard Spain, T.: An assessment of the surface ozone trend in Ireland  
1028 relevant to air pollution and environmental protection, *Atmos Pollut Res*, 3,  
1029 341–351, <https://doi.org/10.5094/APR.2012.038>, 2012.

1030 Tripathi, O. P., Jennings, S. G., Colman, L., Lambkin, K., Moran, E., and Dowd, C. O.: Ozone  
1031 levels , changes and trends over Ireland – an Integrated Analysis, 2013.

1032 Vautard, R., Moran, M. D., Solazzo, E., Gilliam, R. C., Matthias, V., Bianconi, R., Chemel,  
1033 C., Ferreira, J., Geyer, B., Hansen, A. B., Jericevic, A., Prank, M., Segers, A., Silver, J. D.,  
1034 Werhahn, J., Wolke, R., Rao, S. T., and Galmarini, S.: Evaluation of the meteorological  
1035 forcing used for the Air Quality Model Evaluation International Initiative (AQMEII) air  
1036 quality simulations, *Atmos Environ*, 53, 15–37,  
1037 <https://doi.org/10.1016/j.atmosenv.2011.10.065>, 2012

1038 Van Der Werf, G. R., Randerson, J. T., Giglio, L., Collatz, G. J., Mu, M., Kasibhatla, P. S.,  
1039 Morton, D. C., Defries, R. S., Jin, Y., and Van Leeuwen, T. T.: Global fire emissions and the  
1040 contribution of deforestation, savanna, forest, agricultural, and peat fires (1997-2009), *Atmos*  
1041 *Chem Phys*, 10, 11707–11735, <https://doi.org/10.5194/acp-10-11707-2010>, 2010.

1042 Wespes, C., Hurtmans, D., Clerbaux, C., and Coheur, P. F.: O<sub>3</sub> variability in the troposphere  
1043 as observed by IASI over 2008-2016: Contribution of atmospheric chemistry and dynamics,  
1044 *J Geophys Res*, 122, 2429–2451, <https://doi.org/10.1002/2016JD025875>, 2017.

1045 WHO: WHO global air quality guidelines: particulate matter (PM<sub>2.5</sub> and PM<sub>10</sub>), ozone,  
1046 nitrogen dioxide, sulfur dioxide and carbon monoxide, World Health Organization,  
1047 <https://iris.who.int/handle/10665/345329>, 2021.

1048 Wild, O. & Ryan, E. M.: Quantifying and addressing the uncertainties in tropospheric ozone  
1049 and OH in a global chemistry transport model, *EGUsphere* [preprint],  
1050 <https://doi.org/10.5194/egusphere-2025-4534>, 2025.

1051 Yan, Y., et al.: Analysis of European ozone trends in the period 1995–2014, *Atmos. Chem.*  
1052 *Phys.*, 18, 5589–5605, <https://doi.org/10.5194/acp-18-5589-2018>, 2018.

1053 Yerramilli, A., Srinivas Challa, V., Rao Dodla, V. B., Myles, L. T., Pendergrass, W. R., Vogel,  
1054 C. A., Tuluri, F., Baham, J. M., Hughes, R., Patrick, C., Young, J., & Swanier, S. Simulation  
1055 of surface ozone pollution in the Central Gulf Coast region during summer synoptic condition  
1056 using WRF/Chem air quality model. *Atmospheric Pollution Research*, 3(1), 55–71.  
1057 <https://doi.org/10.5094/APR.2012.005>, 2012.

1058 Young, P. J., Archibald, A. T., Bowman, K. W., Lamarque, J., Naik, V., Stevenson, D. S., and  
1059 Tilmes, S.: Pre-industrial to end 21st century projections of tropospheric ozone from the  
1060 Atmospheric Chemistry and Climate Model Intercomparison Project ( ACCMIP ), 2063–  
1061 2090, <https://doi.org/10.5194/acp-13-2063-2013>, 2013.

1062 Yu, H., Chin, M., West, J. J., Atherton, C. S., Bellouin, N., Bergmann, D., Bey, I., Bian, H.,  
1063 Diehl, T., Forberich, O., Hess, P., Schulz, M., Takemura, T., & Tan, Q.. A multimodel  
1064 assessment of the influence of regional anthropogenic emission reductions on aerosol direct  
1065 radiative forcing and the role of intercontinental transport. *Journal of Geophysical Research:*  
1066 *Atmospheres*, 118(13), 7004–7025. <https://doi.org/10.1002/jgrd.50502>,2013.

1067 Zhang, J. J., Wei, Y., & Fang, Z. Ozone pollution: A major health hazard worldwide. *Frontiers*  
1068 *in Immunology*, 10(OCT), 1–10. <https://doi.org/10.3389/fimmu.2019.02518>,2019.

1069 Zhang, C. and Stevenson, D.: Characteristic changes of ozone and its precursors in London  
1070 during COVID-19 lockdown and the ozone surge reason analysis, *Atmos. Environ.*, **273**,  
1071 118980, <https://doi.org/10.1016/j.atmosenv.2022.118980>,2022

1072 Zhu, T., Melamed, M., Parrish, D., Gauss, M., Klenner, L. G., Lawrence, M., Konare, A., and  
1073 Liousse, C.: *Impacts of Megacities on Air Pollution and Climate*, WMO, Geneva, 2012.

1074

1075

1076

1077

1078

## 1079 **Figure Captions**

1080 **Figure 1.** The map of EPA O<sub>3</sub> measurement sites over Ireland with classification of  
1081 backgrounds.

1082 **Figure 2.** Annual average O<sub>3</sub> concentration at different sites in Ireland. In each box, the  
1083 lowest whisker level represents the 5<sup>th</sup> percentile, the box spans from the 25<sup>th</sup> to the 75<sup>th</sup>

1084 percentile, the horizontal line within the box represents the median 50th percentile, and the  
1085 upper whisker represents the 95<sup>th</sup> percentile. The average of monthly O<sub>3</sub> values calculated  
1086 for the entire period of each station, and the red line shows the average monthly O<sub>3</sub>  
1087 variation of all sites top axis shows the month (1– 12).

1088 **Figure 3.** Monthly trend analysis of O<sub>3</sub> at different sites for 10 year period. (2012-2022)  
1089 Adopting the trend reliability scale defined for TOAR-II studies (Chang et al., 2023), trends  
1090 with very high certainty are marked by \*\*\*( $p \leq 0.001$ ), , trends with high certainty with  
1091 \*\*( $p \leq 0.01$ ), and low to medium certainty with \*( $p \leq 0.05$ ). Positive trends are in red  
1092 shade and negative trends are in blue shade.

1093 **Figure 5.-** Trend in O<sub>3</sub> precursors NO<sub>2</sub> (a), and CH<sub>4</sub> (b) at different sites. Trends with very  
1094 high certainty are marked by \*\*\*( $p \leq 0.001$ ), trends with high certainty with \*\*( $p \leq 0.01$ ),  
1095 and low to medium certainty with \*( $p \leq 0.05$ ).

1096 **Figure 6.** Percentage change in NO<sub>2</sub> and O<sub>3</sub> during the lockdown period of 2020 as  
1097 compared to the 2017-2019 average at different sites in Ireland for (a) March (b) April, (c)  
1098 May Month.

1099 **Figure 7.** The comparison of Monthly CAM4 – Chem O<sub>3</sub> and Monthly O<sub>3</sub> observations at  
1100 five sites in Ireland.

1101 **Figure 8.** Absolute contribution of major NO<sub>x</sub> sources (a) (NO<sub>x</sub> Tagging) and VOC source  
1102 (b) (NO<sub>x</sub> Tagging) to the CAM4-Chem simulated surface O<sub>3</sub> for the Mace Head grid cell  
1103 between 2000-2018.

1104 **Figure 9.** Trends in contributions to monthly average modelled Mace Head grid cell surface  
1105 O<sub>3</sub> for the 2000-2018 period derived from (a) NO<sub>x</sub> tagging and (b)VOC tagging.

1106 **Figure 10.** Trend in seasonal Average of observed O<sub>3</sub> (black ) and Model O<sub>3</sub> (red) at Mace  
1107 head, separated into clean sector and EU influenced sector.

1108 **Figure 11.** Exceedances measured at Mace Head per month from 2000 until 2022, during  
1109 the clean air sector(green) and EU influenced sector (yellow). The percentage of both to  
1110 total exceedances is shown in the inlay.

1111 **Figure 12.** (a) The trend in Spring-time exceedances measured at Mace Head between  
1112 2000 and 2022 (blue) with the clean-air exceedances (gold), and (b) The trend in the 95<sup>th</sup>  
1113 percentile of spring (Mar- May) O<sub>3</sub> measured in µg /m<sup>3</sup> for the clean sector (blue) and the  
1114 EU influenced sector (gold).

1115 **Figure 13** Monthly cumulative Mace Head grid cell O<sub>3</sub> contributions to EU influenced  
1116 sector and clean sector exceedances (a) NO<sub>x</sub> tagging and (b) VOC tagging Mace Head grid  
1117 cell.

1118

1119

1120

1121

1122