

1 **Response to Reviewer # 1**

2 We appreciate the anonymous reviewer for the thoughtful reviews and comments. We have carefully
3 considered the suggestions and revised the manuscript accordingly. The **reviewer comments are in blue**,
4 our comments are in black, and **modifications to the manuscript are in red**.

5 **General comments:**

6 Kim et al. investigated the characteristics of organic aerosol (OA) in Seoul during wintertime haze events.
7 In contrast to previous studies conducted in the region, this work successfully identified a nitrogen-
8 containing OA (NOA) factor, likely associated with biomass burning, in addition to five other OA factors
9 using positive matrix factorization (PMF). A notable finding was that the oxidation state of oxygenated OA
10 (OOA) did not align with volatility: the less-oxidized OOA (LO-OOA) was found to be less volatile than
11 the more-oxidized OOA (MO-OOA). This observation challenges the conventional classification of OOA
12 as semi-volatile or low-volatile and supports the use of LO-OOA and MO-OOA terminology.

13 The study provides insights relevant to understanding air quality in Seoul, one of the world's largest
14 metropolitan areas, where haze events are increasingly linked to nitrogen-related pollutants. Although NOA
15 contributed only a minor fraction of total OA, its enhancement offers clues about particle-phase chemistry
16 that may worsen air quality and could inform the design of mitigation strategies. Furthermore, this study
17 represents the first attempt in Seoul to connect the oxidation state of OA with volatility, adding useful
18 information to the literature.

19 However, many of the findings presented are not conceptually new within the broader atmospheric
20 chemistry field. In addition, several interpretations remain speculative, with limited supporting evidence.
21 In particular, Section 3.3 does not convincingly explain why MO-OOA in Seoul appears more volatile than
22 LO-OOA. For these reasons, the work may be more appropriate for publication as a Measurement Report
23 in its current form, unless the discussion and interpretation are substantially strengthened.

24 Thank you for the thoughtful assessment and for recognizing the potential of our study. We appreciate the
25 concern that some interpretations initially appeared speculative. Guided by your comments, we have
26 substantially strengthened Section 3.3 and clarified several points across the manuscript. We believe the
27 revised paper now goes beyond a measurement-only contribution and meets the criteria for a research article
28 in ACP. In addition to reporting winter measurements, the revision explains the chemistry, evaluates
29 alternative explanations, and situates Seoul in comparison with other major cities—consistent with ACP
30 research-article standards rather than a Measurement Report. We believe these advances are appropriate for
31 a research article in ACP.

32 Additional comments are provided below.

33 **Major comments:**

34 Section 3.2: This study is presented as a characterization of wintertime aerosol, yet comparisons are made
35 with fall data. Please clarify the source of the fall dataset—was it obtained from your previous work or from
36 another published study? If it is adopted, this needs to be explicitly stated earlier in the section and
37 consistently noted throughout the manuscript, including in figure captions.

38 Thank you for the comment. Section 3.2 analyzes only the winter-2019 dataset collected in this study. To
39 avoid ambiguity, we clarified the phrasing at the few places where this could be misread.

40 1. **Line 258 :**“Compared to fall ($T_{50} \sim 73$ °C, incomplete evaporation), winter nitrate appeared more
41 volatile, supporting enhanced NOA detection and indicating relatively fewer non-volatile nitrate
42 forms.”**After:** “Compared to previously reported fall conditions ($T_{50} \sim 73$ °C, incomplete

43 evaporation), winter nitrate appeared more volatile, supporting enhanced NOA detection and
44 indicating relatively fewer non-volatile nitrate forms (e.g., Kang et al., 2022; Jeon et al., 2023).”
45 2. **Line 266:** “Residual ammonium at 200 °C was consistent (~4%) across seasons (Kang et al., 2022;
46 Jeon et al., 2023).” **After:** “Residual ammonium at 200 °C was consistent (~4%) in previously
47 reported spring/fall measurements (Kang et al., 2022; Jeon et al., 2023).”
48 3. **Line 268:** “Chloride volatility was also comparable between seasons in terms of T_{50} , but exhibited
49 more complete evaporation in winter (~4% residual vs. ~10% in fall), possibly reflecting a shift in
50 source to more volatile forms like road salt during wintertime.” **After:** “Chloride volatility was also
51 comparable across seasons in prior studies in terms of T_{50} , but exhibited more complete evaporation
52 in winter (~4% residual vs. ~10% in fall), possibly reflecting a shift in source to more volatile forms
53 like road salt during wintertime.”
54 4. Line 274: “This trend aligns with spring and fall observations (Kang et al., 2022; Jeon et al.,
55 2023).” **After:** “This trend aligns with previously reported spring and fall observations (Kang et
56 al., 2022; Jeon et al., 2023).”
57 5. **Line 292:** “This pattern reflects its diverse cooking sources and variable emission profiles (Kang
58 et al., 2022).” **After:** “This pattern reflects its diverse cooking sources and variable emission
59 profiles as previously reported (Kang et al., 2022).”
60 6. **Figure 4 caption :** “Mass fraction remaining (MFR) of non-refractory (NR) aerosol species
61 measured in Seoul using a thermodenuder coupled to a high-resolution time-of-flight aerosol mass
62 spectrometer (HR-ToF-AMS). Winter 2019 (this study; dashed) is compared with fall 2019
63 (previously reported; solid) (Jeon et al., 2023). Species include organics (magenta), nitrate (blue),
64 sulfate (orange), ammonium (green), and chloride (red).

65 Section 3.3: The discussion in this section could be substantially strengthened. For instance, the relatively
66 high volatility of MO-OOA may represent a distinct feature of Seoul compared to other megacities. A more
67 thorough comparison with results from other TD-AMS studies would help contextualize your findings.
68 Given that one of the stated objectives of this work is to “improve understanding of wintertime OA in Seoul,”
69 highlighting how Seoul’s OOA differs from or aligns with other urban environments would significantly
70 enhance the discussion.

71 Thank you. Per your suggestion we’ve strengthened section 3.3 with explicit TD-AMS comparisons to
72 other megacities in the discussion of section 3.3. The updated paragraph (below) now cites representative
73 studies from Mexico City/Los Angeles, Paris, Beijing, and Shenzhen, plus springtime Seoul for seasonal
74 context.

75 “Viewed against prior TD-AMS results, the volatility of Seoul’s winter MO-OOA presents a unique case,
76 particularly in the nature of its O:C-volatility relationship. Prior urban studies have commonly reported
77 substantial SVOC-OA, consistent with high photochemical activity or elevated loadings; for example,
78 Mexico City/Los Angeles showed pronounced SVOC-IVOC contributions during warm seasons (Cappa
79 and Jimenez, 2010), and summertime Beijing and wintertime Shenzhen likewise exhibited strong overall
80 OA volatility (Xu et al., 2019; Cao et al., 2018). While these comparisons establish that volatile OA is
81 common, they generally did not report the factor-level inversion observed here, where the highly-oxidized
82 OOA component (MO-OOA) was *more* volatile than a less-oxidized OOA (LO-OOA). This behavior is
83 distinct from findings in colder, lower-loading regimes; wintertime Paris, for instance, maintained the
84 conventional hierarchy where the more-oxidized OOA was comparatively less volatile (Paciga et al., 2016).
85 Furthermore, seasonal context within Seoul showed springtime OA with lower oxidation levels than our
86 winter MO-OOA despite similar SVOC contributions (Kang et al., 2022). This comprehensive comparison
87 underscores the unusual nature of the O:C-volatility relationship observed under the specific winter haze
88 conditions in Seoul.

89 Line 163: The claim that enhanced MO-OOA and NOA formation during haze is due to stagnation is not
90 convincing without supporting meteorological data. Please provide wind speed/direction or back-trajectory
91 analyses.

92 Thank you for this comment. We now document meteorological support for stagnation during haze by (i)
93 presenting back-trajectory clustering that identifies a local recirculation regime (Cluster 1; Fig. S8) and (ii)
94 showing persistently low wind speeds during haze windows in the campaign time series (Fig. S6). These
95 diagnostics indicate that the haze episodes were dominated by local accumulation rather than long-range
96 transport, consistent with the observed co-enhancement of MO-OOA and NOA. Now the relevant section
97 reads;

98 **“Back-trajectory clustering shows frequent short-range recirculation over the Seoul Metropolitan Area
99 during haze (Cluster 1; Fig. S8), and the time series indicates persistently low surface wind speeds during
100 these periods (1.73 ± 0.89 vs. 2.34 ± 1.18 (clean)) (Fig. S6), together pointing to stagnation-driven
101 accumulation of local emissions; the concurrent increases in MO-OOA and NOA are therefore consistent
102 with enhanced in-city formation under stagnant conditions.”**

103 Line 172: Could factor resolution differ between V- and W-mode? Please show whether NOA is detected
104 in both modes and perform PMF sensitivity tests to mode/resolution.

105 Thank you for the careful question. In the HR-ToF-AMS, W-mode provides higher mass resolving power
106 than V-mode, enabling separation of near-isobaric ions (e.g., m/z 58, 86, 100 series) that are important for
107 resolving amine-rich (NOA) spectra; V-mode offers higher sensitivity but lower resolution and is widely
108 used for quantification and bulk metrics, whereas W-mode is preferred for factor analysis of HR ion families.
109 Consistent with standard AMS practice, our PMF was performed on W-mode HR spectra (Methods), where
110 the amine fragments are better separated and uncertainties are well characterized. Running PMF on the
111 lower-resolution V-mode would merge key N-containing ions with interferences and degrade factor
112 identifiability, so we did not perform a V-mode PMF.

113 Line 194: Since amines can also originate from solvents, might some portion of NOA be linked to solvent
114 emissions? Given Seoul’s solvent usage, this possibility should be addressed.

115 Thank you for the helpful comment. We agree that solvent- and industry-related amines (e.g., tertiary
116 amines used in polyurethane manufacturing/foams, metalworking, printing, and wastewater treatment) can
117 contribute to urban amine budgets. Reviews and technical sources document such emissions and uses (e.g.,
118 TEA/DEEA/DMEA in PU catalysts), and Korean inventories indicate that solvent use is a major component
119 of anthropogenic VOC emissions nationally (references). However, In our dataset, several diagnostics favor
120 near-source combustion/BB influences for NOA during winter in Seoul: (i) night–early-morning peaks and
121 co-variation with BBOA, (ii) back-trajectory clusters indicating local recirculation, and (iii) mass-spectral
122 features dominated by small alkylamine fragments (e.g., m/z 30, 44, 58, 86) that match low-molecular-
123 weight alkylamines commonly associated with combustion/BB rather than tertiary amines typical of
124 PU/solvent use. That said, we now explicitly note that a minor contribution from solvent/industrial amines
125 cannot be ruled out, and we added references and a sentence in the Discussion to reflect this possibility.

126

127 **“These amines are commonly emitted during the combustion of nitrogen-rich biomass and proteinaceous
128 materials and are frequently associated with biomass-burning emissions (Ge et al., 2011). Previous
129 molecular analyses in Seoul also indicate DMA, MA, and TMA as the dominant amine species in December
130 (Baek et al., 2022). While other amines such as triethylamine (TEA), diethylamine (DEA), and ethylamine**

131 (EA) may contribute via industrial/solvent pathways (e.g., chemical manufacturing, petrochemical
132 corridors, wastewater treatment), our HR-AMS spectra are dominated by small alkylamine fragments (*m/z*
133 30, 44, 58, 86) and the diurnal behavior co-varies with combustion markers (below), indicating a primarily
134 combustion-linked influence. Nevertheless, recent urban measurements and sector-based analyses show
135 that industrial activities can contribute measurable amines in cities (Tiszenkel et al., 2024; Zheng et al.,
136 2015; Mao et al., 2018; Shen et al., 2017; Yao et al., 2016). Accordingly, a minor NOA contribution from
137 solvent/industrial amines cannot be excluded.”

138

139 Line 199: Where do you expect biomass burning that could influence Seoul’s air to occur? Since NOA
140 (linked to BBOA) is a central finding, more discussion on plausible sources and transport pathways would
141 be valuable.

142 Thank you for this valuable suggestion. We expanded the discussion to identify plausible biomass-burning
143 (BB) source regions and pathways that could influence Seoul and to clarify why our dataset points primarily
144 to local influences. Specifically: (i) wintertime urban/peri-urban small-scale BB (residential solid-fuel use,
145 restaurant charcoal use, and intermittent waste burning) has been observed in Seoul and surroundings
146 (Kim et al., 2017); (ii) agricultural residue burning occurs seasonally in nearby provinces and can
147 episodically affect the Seoul Metropolitan Area (Han et al., 2022); and (iii) regional transport from upwind
148 regions (including northeastern China/North Korea) can influence Korea under northerly/northwesterly
149 flow (Lamb et al., 2018; Nault et al., 2018). In our case, the night–early-morning peaks, co-variation with
150 BBOA, and back-trajectory clusters indicating local recirculation are most consistent with near-source
151 emissions within the metropolitan area during the study period (Yoo et al., 2024). We added further
152 discussion at the end of section 3.1.3.

153 "Regarding source location, several pathways can influence Seoul’s biomass burning signature. First,
154 urban/peri-urban small-scale burning (e.g., solid-fuel use in select households, restaurant charcoal use, and
155 intermittent waste burning) has been reported and can enhance BBOA locally (Kim et al., 2017). Second,
156 nearby agricultural-residue burning in surrounding provinces occurs seasonally and can episodically impact
157 the metropolitan area (Han et al., 2022). Third, regional transport from upwind regions (e.g., northeastern
158 China/North Korea) can bring biomass burning influenced air masses under northerly/northwesterly flow
159 (Lamb et al., 2018; Nault et al., 2018). In this dataset, the nighttime and early-morning enhancements, the
160 BBOA–NOA co-variation, and trajectory clusters showing regional recirculation indicate a predominantly
161 local/near-source contribution during the study period (Yoo et al., 2024), with episodic non-local influences
162 remaining possible."

163

164 Line 225: The phrase “...with inorganic secondary species such as biomass burning” is unclear. Please
165 revise.

166 Thank you for pointing this out. We agree the phrase was ambiguous—biomass burning is a source, not an
167 inorganic secondary species. We removed the confusing wording and clarified the comparison to sulfate,
168 nitrate, and ammonium. Now the relevant line reads:

169 “In contrast, LO-OOA exhibited only modest correlations with sulfate, nitrate, and ammonium ($r = 0.50$,
170 0.51, and 0.42, respectively), suggesting additional contributions from semi-primary sources not closely
171 linked to inorganic secondary formation (e.g., cooking, traffic, biomass burning).”

172 Line 226: Earlier you noted that LO-OOA lacks m/z 60, but here you attribute it partly to combustion-
173 related activities. This appears contradictory. Please clarify or rewrite.

174 Thank you for noting the ambiguity. We agree that invoking *m/z* 60 (levoglucosan) alongside an earlier
175 statement that LO-OOA lacks *m/z* 60 is confusing. Now the section reads:

176 “LO-OOA does not exhibit a pronounced *m/z* 60 (levoglucosan) signal (Fig. S2); however, the
177 levoglucosan marker (f_{60}) is known to diminish with atmospheric aging and can become weak or
178 undetectable downwind (Hennigan et al., 2010; Cubison et al., 2011). Taken together, the weaker coupling
179 with secondary inorganics and the absence of a strong *m/z* 60 peak indicate that LO-OOA is a mixture of
180 aged secondary organics and semi-primary urban emissions, while a contribution from aged biomass-
181 burning influence cannot be ruled out.”

182 Line 257–258: The text seems to imply that NOA could be detected as nitrate. However, most reduced
183 nitrogen species (other than nitro-aromatics, based on Xu et al., 2021, AMT) are unlikely to be detected as
184 nitrate by AMS. Please clarify.

185 Thank you for pointing this out. We agree that reduced-nitrogen species (e.g., amines) are not detected by
186 the AMS as “nitrate.” In the AMS, the nitrate channel reflects inorganic nitrate and organic nitrates (RONO_2)
187 via $\text{NO}^+/\text{NO}_2^+$ fragments; most reduced-N compounds do not contribute, with the noted exception that nitro-
188 aromatics can appear in the nitrate channel. To avoid confusion, we removed NOA from this sentence and
189 now discuss nitrate volatility independently.

190 Compared to previously reported fall conditions ($T_{50} \sim 73$ °C, incomplete evaporation), winter nitrate
191 appeared more volatile, indicating relatively fewer non-volatile nitrate forms (e.g., Kang et al., 2022; Jeon
192 et al., 2023).

193 Line 261: If metallic sulfates contribute to PM_1 , then defining PM_1 as $\text{NR-PM}_1 + \text{BC}$ may underestimate
194 total mass. Please comment. Thank you for the helpful comment. Metallic sulfates (e.g., CaSO_4 , MgSO_4 ,
195 Na_2SO_4) would not be fully captured by the AMS and, if abundant in PM_1 , could bias $\text{PM}_1 = \text{NR-PM}_1 + \text{BC}$
196 low. However, for our winter urban period in Seoul:

- 197 1. The HR-ToF-AMS measures non-refractory submicron species; refractory salts typically do not
198 flash-vaporize at the AMS vaporizer and contribute negligibly to the AMS sulfate signal (Allan et
199 al., 2004; Canagaratna et al., 2007).
- 200 2. Multiple Seoul/Korea studies show PM_1 is dominated by NR species + BC, with sea-salt/crustal
201 components largely in the coarse mode; our analysis window also excluded dust events, so
202 refractory PM_1 was minimal .
- 203 3. The subtle ~140 °C slope change is more plausibly due to ammonium-sulfate system behavior (e.g.,
204 morphology/phase-state changes in $(\text{NH}_4)_2\text{SO}_4/\text{NH}_4\text{HSO}_4$ and mixed organic–inorganic particles)
205 and/or organosulfate mixing, rather than metallic sulfates (Huffman et al., 2009; Faulhaber et al.,
206 2009).

207 Accordingly, using $\text{PM}_1 = \text{NR-PM}_1 + \text{BC}$ does not materially underestimate PM_1 for this dataset. We added
208 a sentence in Methods clarifying the non-refractory definition and a note in Results that we interpret the
209 ~140 °C feature as ammonium-sulfate/phase-state and/or organosulfate mixing—not metallic sulfate. Now
210 the relevant section reads;

211 Line 261: “Sulfate was the least volatile ($T_{50} \approx 170$ °C), consistent with ammonium sulfate (Scott and Cattell,
212 1979). A subtle slope change near 140 °C likely reflects ammonium-sulfate morphology/phase-state

213 changes and/or organosulfate–inorganic mixing, rather than contributions from metallic (refractory)
214 sulfates, which are not efficiently detected by AMS.”

215 Methods :“PM₁ mass in this study is taken as NR-PM₁ (from AMS) + black carbon (BC; measured by
216 MAAP), which is appropriate for winter Seoul where refractory PM₁ (metal/sea-salt/crustal) is minor and
217 dust events were excluded (e.g., Kim et al., 2017; Nault et al., 2018; Kang et al., 2022; Jeon et al., 2023).”

218 Line 289: Since MO-OOA is described as more volatile, the phrase “consistent with its aged, highly
219 condensed nature” seems contradictory. Consider deleting or rephrasing.

220 Thank you for the suggestion. We deleted the phrase.

221 Line 308: What alternative oxidation mechanisms or environmental conditions could explain the observed
222 inverse relationship between O:C ratio and volatility? Please specify and cite relevant studies.

223 Thank you for the comment. We agree the original lead-in was too vague on the mechanisms. We revised
224 the sentence to explicitly foreshadow the explanation that immediately follows (autoxidation-driven
225 fragmentation under low-OH/low-O₃ haze conditions and enhanced condensation at high particle loadings).
226 We also added a section pointer for clarity.

227 “This observation appears to contradict the usual inverse O:C–volatility relationship; however, under winter
228 haze conditions—with suppressed O₃/low OH, particle-phase autoxidation and fragmentation can yield
229 higher-O:C yet more volatile products, with enhanced condensation on abundant particle surface area
230 (details below).”

231 Line 311: The authors attribute reduced OH to suppressed O₃ photolysis under haze. However, in polluted
232 boundary-layer conditions, OH production often depends strongly on HONO photolysis in addition to O₃
233 photolysis. If the authors want to keep this statement, I suggest add discussion and relevant reference on
234 this.

235 Thank you for the comment. We did not measure HONO directly, so our “low-OH” interpretation was
236 based primarily on reduced O₃ photolysis during haze. We now clarify that haze generally reduces actinic
237 flux for all photolysis channels, including HONO photolysis, due to aerosol extinction/scattering. At the
238 same time, HONO concentrations often increase at night during polluted, humid conditions because of
239 heterogeneous NO₂ → HONO production on surfaces/aerosols and direct/near-surface emissions; thus the
240 net OH from HONO can remain significant or even episodically important despite lower *j*(HONO). We
241 added text noting both effects and cited field studies in Korea (KORUS-AQ) and China that document (i)
242 reduced photolysis frequencies under haze and (ii) enhanced HONO production via heterogeneous
243 pathways. This balanced perspective does not change our main conclusion (a low-OH regime relative to
244 sunny, high-photolysis periods), but it acknowledges the potential compensating role of HONO under haze.

245 “We note that haze also suppresses HONO photolysis; however, HONO concentrations can be elevated at
246 night and early morning via heterogeneous NO₂ conversion and surface emissions, so net OH from HONO
247 may remain non-negligible even as photolysis rates are depressed (e.g., Gil et al., 2021; Kim et al., 2024;
248 Slater et al., 2020).

249 Line 312: Under conditions of reduced OH, how are RO₂ radicals formed and how might they contribute to
250 particle-phase autoxidation? Please elaborate and provide references.

251 Thank you for the prompt. We clarified that even under reduced OH (low $j(O^1D)$ during haze), RO_2 can
252 arise from (i) residual gas-phase ozonolysis of alkenes (Ehn et al., 2014; Ziemann and Atkinson, 2012;
253 Bianchi et al., 2019), (ii) nighttime/low-light NO_3 -initiated chemistry where NO_x and remaining O_3 permit
254 NO_3 formation (Brown and Stutz, 2012; Ziemann and Atkinson, 2012), and (iii) condensed-phase radical
255 chemistry/autoxidation in particles once peroxy radicals are initiated (Kroll and Seinfeld, 2015; Berndt et
256 al., 2016; Bianchi et al., 2019). These routes can sustain RO_2 pools that drive particle-phase autoxidation
257 under haze. Now the line reads;

258 “Even under low-OH conditions, NO_3 formed via $NO_2 + O_3$ can initiate RO_2 production through addition
259 to alkenes, while reduced photolysis at night/low light extends NO_3 lifetimes; these RO_2 then participate in
260 particle-phase autoxidation, yielding highly oxygenated yet relatively volatile products (Brown and Stutz,
261 2012; Ziemann and Atkinson, 2012; Ehn et al., 2014; Berndt et al., 2016; Bianchi et al., 2019).”

262 Line 315–316: To support the interpretation that fragmentation contributes to the high volatility of MO-
263 OOA, please cite previous studies that compared OA composition using both online and offline techniques.

264 Thank you for the comments. We now cite studies that combine online (AMS/FIGAERO-CIMS) and offline
265 (filter HRMS/GC-MS) techniques showing that fragmentation-dominated aging yields high-O:C yet more
266 volatile products and elevated f_{44} without proportional increases in f_{43} .

267 “Consistent with this interpretation, online AMS/FIGAERO-CIMS and EESI-TOF, as well as offline
268 HRMS/GC-MS, have reported high-O:C yet more-volatile product distributions accompanied by elevated
269 f_{44} with comparatively stable f_{43} under fragmentation-dominated aging (Kroll et al., 2009; Ng et al., 2010;
270 Chhabra et al., 2011; Lambe et al., 2012; Lopez-Hilfiker et al., 2016; D’Ambro et al., 2017).”

271 Line 317: Please revise “semi-volatile species” to “semi-volatile/intermediate-volatility organics.”
272 Additionally, please clearly comment that functionalized but low-molecular-weight compounds can fall in
273 the SVOC–IVOC range and may contribute to the high volatility of MO-OOA. Please provide supporting
274 references.

275 We revised “semi-volatile species” to “semi-volatile/intermediate-volatility organics (SVOC–IVOC)” and
276 explicitly note that functionalized, low-molecular-weight compounds can fall in the SVOC–IVOC range
277 and thus contribute to the high apparent volatility of MO-OOA. References added.

278 “Furthermore, high aerosol mass loadings during haze events provide abundant surface area for the uptake
279 of semi-volatile/intermediate-volatility organics (SVOC–IVOC) via absorptive partitioning, so that higher
280 Co_a enhances condensation (Pankow, 1994; Donahue et al., 2006; Hallquist et al., 2009; Robinson et al.,
281 2007). We also note that functionalized, low-molecular-weight compounds can reside in the SVOC–IVOC
282 range and thus contribute to the high apparent volatility of MO-OOA (Ng et al., 2010; Chhabra et al., 2011;
283 Lopez-Hilfiker et al., 2016; D’Ambro et al., 2017).”

284

285 Line 327–328: A comparison of average OA mass spectra between haze and non-haze periods would be
286 informative. Do non-haze periods show relatively more high-molecular-weight fragments?

287 Thank you for the insightful comment. We added a direct comparison of average OA mass spectra for haze
288 and non-haze periods (new Fig. S12; events defined in Fig. S6). Because AMS EI at 70 eV produces
289 extensive fragmentation, non-haze periods do **not** exhibit “more high-molecular-weight fragments.” The
290 differences are expressed in low- m/z markers: haze spectra show enhanced oxygenated fragments (m/z 28

291 = CO⁺, 29 = CHO⁺, 44 = CO₂⁺) together with higher f_{44} and O:C, while f_{43} remains comparatively unchanged.
292 Non-haze spectra display larger fractional contributions of hydrocarbon-like ions (*m/z* 41, 43, 55, 57) and
293 lower f_{44} and O:C. This pattern—increased f_{44} with flat f_{43} —is the canonical signature of fragmentation-
294 dominated aging rather than functionalization (Kroll et al., 2009), reinforcing our interpretation that the
295 periods of elevated MO-OOA reflect advanced oxidation and increased volatility arising from autoxidation
296 and the condensation of small oxygenated fragments. Now the sentence reads;

297 “Consistent with this, the haze–non-haze comparison, including the high-MO-OOA interval (Fig. S12),
298 shows larger oxygenated fragments (*m/z* 28, 29, 44) and higher f_{44} and O:C during haze, whereas non-haze
299 periods exhibit relatively larger fractional hydrocarbon fragments (*m/z* 41, 43, 55, 57). These spectral
300 contrasts indicate that the elevated volatility of MO-OOA reflects advanced oxidation—via autoxidation
301 and the condensation of small oxygenated fragments—rather than enrichment of high-molecular-weight
302 ions, particularly under conditions of limited OH and high particulate surface area.”

303 Specific comments:

304 2 Experimental methods

305 Line 97: Since you are presenting quantitative results, please specify what collection efficiency (CE) value
306 was applied in the AMS data analysis. In addition, could the use of the thermal denuder (TD) influence CE,
307 for example by altering particle phase state or mixing characteristics? Please clarify whether you assumed
308 the same CE with and without TD and provide justification or relevant references.

309 Thank you for this point. We apply time-resolved, composition-dependent CE(t) (Middlebrook et al., 2012)
310 separately to the bypass and TD lines. This approach allows CE to vary with the evolving nitrate fraction
311 and neutralization, which TD heating can influence. The campaign-average CE values were TD: 0.55 ±
312 0.08 and bypass: 0.53 ± 0.04; the mean difference is 0.02 (~3.7%), which is smaller than the combined
313 uncertainty of the two estimates (~0.09). In other words, the CE(t) distributions are similar between lines
314 (comparable means and standard deviations), and we did not force them to be equal—each line’s CE(t) was
315 used in its own mass quantification.

316 We agree that phase-state or mixing changes beyond composition could, in principle, perturb CE. However,
317 prior TD–AMS studies report that such TD-induced CE effects are modest and predominantly
318 multiplicative/near-uniform with respect to thermogram interpretation, not altering thermogram shapes or
319 T₅₀ ordering (Huffman et al., 2009; Faulhaber et al., 2009; Cappa & Jimenez, 2010). Given our use of CE(t)
320 that already responds to composition changes, the small TD–bypass CE difference relative to uncertainty,
321 and literature showing minor additional CE impacts from phase/bounce, we consider any residual CE
322 effects beyond composition to be minor for the volatility metrics presented. We have clarified this in the
323 Methods and added the supporting citations.

324 “NR-PM₁ quantification followed established AMS protocols (Ulbrich et al., 2009; Zhang et al., 2011).
325 Both the bypass and TD streams were processed using a time-resolved, composition-dependent collection
326 efficiency CE(t) following Middlebrook et al. (2012). TD heating can modify particle water and phase
327 state/mixing and thereby influence CE beyond composition (Huffman et al., 2009), but prior TD–AMS
328 studies indicate that such effects are modest and largely multiplicative, which do not distort thermogram
329 shapes or T₅₀ ordering (Faulhaber et al., 2009; Cappa & Jimenez, 2010). In our data, the CE(t) statistics for
330 the two lines were similar (campaign-average CE: TD = 0.55 ± 0.08; bypass = 0.53 ± 0.04; Δ = 0.02 ≈ 3.7%,
331 below the combined uncertainty ≈ 0.09). We therefore report volatility metrics with these line-specific CE(t)
332 corrections applied and interpret potential residual CE effects as minor.”

333 Finally, if you define PM_1 mass as NR- PM_1 + BC, it would be important to cite studies demonstrating that
334 PM_1 in Seoul contains only limited amounts of metals or other refractory components (e.g., dust, sea salt),
335 otherwise this definition may underestimate the total PM_1 mass.

336 Thank you for raising this point. In Seoul, multiple ambient studies indicate that submicron mass (PM_1) is
337 dominated by non-refractory species (organics, sulfate, nitrate, ammonium) plus BC, with refractory
338 metal, dust and sea salt contributing only a small fraction to PM_1 (and mainly affecting supermicron/coarse
339 modes). Our study period excludes dust events. Accordingly, defining $PM_1 = NR-PM_1 + BC$ is consistent
340 with prior Seoul observations and with regional aircraft/surface measurements during KORUS-AQ. We
341 have added the following references and a brief note in Methods to clarify this assumption.

342 “ PM_1 mass in this study is taken as NR- PM_1 (from AMS) + black carbon (BC; measured by MAAP), which
343 is appropriate for winter Seoul where refractory PM_1 (metal/sea-salt/crustal) is minor and dust events were
344 excluded (e.g., Kim et al., 2017; Nault et al., 2018; Kang et al., 2022; Jeon et al., 2023)..”

345 [3 Results and discussion](#)

346 [Line 149: Please clarify why a PM level of \$\sim 28 \mu\text{g m}^{-3}\$ is considered “moderate.” A reference or comparison](#)
347 [to regional air quality standards would help.](#)

348 Thank you for the suggestion. We now justify “moderate” using (i) prior winter PM_1 in Seoul ($27.5 \mu\text{g m}^{-3}$
349 on average; stagnant episodes $\approx 44 \mu\text{g m}^{-3}$), showing our mean ($27.8 \pm 15.3 \mu\text{g m}^{-3}$) is comparable to the
350 historical winter mean and below stagnant loads, and (ii) a Korea-specific $PM_1/PM_{2.5} \approx 0.8$, which implies
351 an equivalent $PM_{2.5} \approx 34.8 \mu\text{g m}^{-3}$, near the national 24-h $PM_{2.5}$ standard ($35 \mu\text{g m}^{-3}$). We have added these
352 references and clarifications.

353

354 “[We conducted continuous measurements from 28 November to 28 December 2019, characterizing a winter](#)
355 [period with a mean \$PM_1\$ concentration of \$27.8 \pm 15.3 \mu\text{g m}^{-3}\$. This concentration is characterized as moderate;](#)
356 [it closely matches historical winter \$PM_1\$ means in Seoul \(Kim et al., 2017\) and implies an equivalent \$PM_{2.5}\$](#)
357 [concentration is about \$34.8 \mu\text{g m}^{-3}\$ \(using a Korea-specific \$PM_1/PM_{2.5} \approx 0.8\$ \(Kwon et al., 2023\), which is](#)
358 [near the national 24-h \$PM_{2.5}\$ standard \(\$35 \mu\text{g m}^{-3}\$ \) \(AirKorea\). ”](#)

359 [Line 186: A reference regarding the atmospheric lifetime and reactivity of amines would strengthen this](#)
360 [discussion.](#)

361 Thank you for the helpful suggestion. We added citations documenting the short atmospheric lifetimes and
362 high reactivity of alkyl amines due to fast gas-phase oxidation .

363 “[Back-trajectory analysis linked these events to regional recirculation patterns \(Cluster 1, Fig. S7\),](#)
364 [suggesting a predominantly local origin—consistent with the short atmospheric lifetimes and high reactivity](#)
365 [of most amines \(Ge et al., 2011; Nielsen et al., 2012; You et al., 2014; Hanson et al., 2011\). ”](#)

366 [Line 242: Please italicize all “ \$m/z\$ ##” and “ \$f##” notations throughout the manuscript for consistency.\$](#)

367 Thank you for pointing this out. We have applied consistent typography throughout the manuscript by
368 italicizing all m/z and f notations (e.g., m/z 44, f 44). This change has been implemented across the main text,
369 figure and table captions, and the Supplementary Information to ensure uniform style.

370 [Line 246: Instead of describing the correlation between NOA and BBOA as “moderate,” please quantify](#)
371 [the correlation coefficient and state whether it is higher relative to other factor pairs.](#)

372 Thank you for the suggestion. We replaced the qualitative phrase “moderate” with quantitative values and
373 clarified the context relative to other factor pairs. Specifically, the BBOA–NOA time-series correlation is
374 **r = 0.30** (the highest among the NOA–factor pairs examined), and their diurnal-profile correlation is **r = 0.63**. We retain the ion-level correlations for completeness.

376 “Quantitatively, BBOA and NOA correlate with $r = 0.30$ in the time series—the highest among the NOA–
377 factor pairs in our dataset—and their diurnal profiles are more strongly correlated ($r = 0.63$). Consistent
378 with this, nitrogen-containing ions characteristic of NOA, $\text{C}_2\text{H}_4\text{N}^+$ ($r = 0.67$) and $\text{C}_2\text{H}_6\text{N}^+$ ($r = 0.56$), co-vary
379 with BBOA and are dominant peaks in the NOA mass spectrum (Fig. 2; Fig. S3; Fig. S10; Fig. S11).”

380 Line 248: The phrase “such as biomass burning” may be redundant here, since this point is already discussed
381 in the BBOA section. Consider removing it.

382 Thank you for the suggestion. We agree that “such as biomass burning” is redundant given the discussion
383 in the BBOA section. We have removed the phrase and streamlined the sentence for clarity. The sentence
384 now reads:

385 “This overlap suggests a potential shared emission source or co-emission scenario, consistent with the co-
386 emission of both organic aerosols and reduced nitrogen-containing compounds”

387 Line 300–303: Please provide supporting references for this discussion.

388 We added references demonstrating that functionalization typically lowers volatility, while fragmentation
389 increases it and can reduce SOA yields despite higher O:C; this supports our emphasis on pairing
390 composition with direct volatility constraints. Now the section reads;

391 “the addition of oxygen-containing functional groups (e.g., hydroxyl, carboxyl, carbonyl) increases
392 molecular weight and enhances intermolecular interactions such as hydrogen bonding, thereby reducing
393 vapor pressure (Jimenez et al., 2009; Kroll and Seinfeld, 2008). Moreover, oxidative aging often leads to
394 oligomerization or functionalization, promoting particle-phase retention and reducing the effective
395 **saturation concentration (C*)** (Donahue et al., 2011; Robinson et al., 2007).”

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579 **Response to Reviewer # 2**

580 We appreciate the anonymous reviewer for the thoughtful reviews and comments. We have carefully
581 considered the suggestions and revised the manuscript accordingly. The **reviewer comments are in blue**,
582 our comments are in black, and **modifications to the manuscript are in red**.

583 Kim et al. present a study of aerosol chemistry during wintertime in Seoul in 2019. Building upon their
584 earlier publication in 2017, the authors conducted a dedicated field campaign focusing on aerosol volatility
585 and obtained several intriguing results that are highly relevant to understanding and modeling aerosol
586 processes in this megacity. One particularly interesting finding is that highly oxidized organic aerosols were
587 shown to be highly volatile, providing observational evidence for autoxidation and fragmentation processes
588 occurring in the particle phase. The dataset is well analyzed, and the manuscript is clearly written. The
589 study fits well within the scope of ACP, and I consider it suitable for publication as a research article, rather
590 than a measurement.

591 We appreciate the thoughtful reviews and comments. We appreciate the recognition that our findings,
592 particularly the observational evidence for autoxidation and fragmentation leading to highly oxidized but
593 volatile organic aerosols, are highly relevant to understanding and modeling aerosol processes in the region.
594 We have carefully considered the suggestions and revised the manuscript accordingly. The **reviewer**
595 **comments are in blue**, our comments are in black, and **modifications to the manuscript are in red**.

596 **Section Introduction:** The Introduction could be further strengthened by expanding the background on
597 aerosol volatility. Including more context on previous volatility-related studies would help frame the
598 contribution of this work.

599 Thank you for this helpful suggestion. To clarify the aims and frame our contribution, we expanded the
600 Introduction to explain why OA volatility is central to gas–particle partitioning and model performance,
601 and summarize frameworks that couple volatility and oxidation. These changes/additions help motivate our
602 TD-AMS approach and the winter focus of this study. Now the relevant section reads:

603 “Volatility is a key parameter for characterizing organic aerosol (OA) properties, as it governs gas-to-
604 particle partitioning behavior and directly influences particle formation yields (Sinha et al., 2023). The
605 classification of OA species based on their volatility—from extremely low-volatility (ELVOC) to semi-
606 volatile (SVOC) and intermediate-volatility (IVOC) compounds—is central to the conceptual framework
607 of secondary OA (SOA) formation and growth (Donahue et al., 2006). It also affects atmospheric lifetimes
608 and human exposure by determining how long aerosols remain suspended in the atmosphere (Glasius and
609 Goldstein, 2016). Therefore, accurately capturing OA volatility is essential for improving predictions of
610 OA concentrations and their environmental and health impacts. However, chemical transport models often
611 significantly underestimate OA mass compared to observations (Matsui et al., 2009; Jiang et al., 2012; Li
612 et al., 2017), largely due to incomplete precursor inventories and simplified treatment of processes affecting
613 OA volatility. For instance, aging—through oxidation reactions such as functionalization and
614 fragmentation—can significantly alter volatility by changing OA chemical structure (Robinson et al., 2007;
615 Zhao et al., 2016). Early volatility studies primarily utilized thermal denuders (TD) coupled with various
616 detection instruments to investigate the thermal properties of bulk OA (Huffman et al., 2008). The
617 subsequent coupling of TD with the Aerosol Mass Spectrometer allowed for component-resolved volatility

618 measurements, providing critical, quantitative insight into the properties of OA factors (e.g., SV-OOA vs.
619 LV-OOA) across different regions (Paciga et al., 2016; Cappa and Jimenez, 2010). These component-
620 resolved volatility data are often used to constrain the Volatility Basis Set (VBS)—the current state-of-the-
621 art framework for modeling OA partitioning and evolution (Donahue et al., 2006). However, a limitation
622 in many field studies is that the TD-AMS thermogram data are rarely translated into quantitative VBS
623 distributions for individual OA factors, which limits their direct use in chemical transport models.
624 Furthermore, the volatility of OOA during extreme haze conditions, where the expected inverse correlation
625 between oxidation (O:C) and volatility can break down (Jimenez et al., 2009), remains poorly characterized,
626 particularly in East Asia's highly polluted winter environments. A recent study in Korea further highlighted
627 the importance of accounting for such processes when interpreting OA volatility under ambient conditions
628 (Kang et al., 2023). Given its central role in OA formation, reaction, and atmospheric persistence, volatility
629 analysis is critical for bridging the gap between measurements and model performance.”

630 Section 2.1: This section currently begins with a citation to the authors' earlier study, which may lead
631 readers to assume that the dataset is the same. However, the actual description of the 2019 field study only
632 appears several sentences later. To improve clarity and logical flow, I suggest first presenting the details of
633 the current field campaign and then referring back to the earlier study for context.

634 We appreciate the suggestion. We rewrote the opening of section 2.1 to first describe the 2019 field
635 campaign and site, followed by a brief pointer to our earlier winter study for background on the same
636 location. We believe this improves the logical flow and avoids confusion with the 2017 dataset. Now the
637 relevant section reads;

638 “We conducted continuous real-time measurements in Seoul, South Korea, from 28 November to 28
639 December 2019. The sampling site was located in the northeastern part of the city (37.60° N, 127.05° E),
640 approximately 7 km from the city center, surrounded by major roadways and mixed commercial–residential
641 land use. Air samples were collected at an elevation of approximately 60 meters above sea level, on the
642 fifth floor of a building. A detailed site description has been reported previously for winter Seoul (Kim et
643 al., 2017). During this period, the average ambient temperature was 1.76 ± 4.3 °C, and the average relative
644 humidity (RH) was $56.9 \pm 17.5\%$, based on data from the Korea Meteorological Administration
645 (<http://www.kma.go.kr>).”

646

647 Line 146-150: This paragraph relies on information from the Supplementary Material, which makes it
648 awkward as an entry point into the main results. I suggest either removing it or integrating the content later
649 in the manuscript, once the main results are introduced.

650 Thank you for the helpful suggestion. We agree that starting a section by referencing time series data in the
651 Supplementary Material (Fig. S6) made the initial description awkward. We have reorganized the opening
652 of the Results section (section 3.1). The revised introductory paragraph now immediately establishes the
653 campaign's characteristics, providing the necessary meteorological and concentration context before
654 detailing the composition in Fig. 1. The revised section now reads:

655 “We conducted continuous measurements from 28 November to 28 December 2019, characterizing a winter
656 period with a mean PM_1 concentration of $27.8 \pm 15.3 \mu\text{gm}^{-3}$. This concentration is characterized as moderate;
657 it closely matches historical winter PM_1 means in Seoul (Kim et al., 2017) and implies an equivalent $\text{PM}_{2.5}$
658 concentration is about $34.8 \mu\text{gm}^{-3}$ (using a Korea-specific $\text{PM}_1/\text{PM}_{2.5} \approx 0.8$ (Kwon et al., 2023)), which is
659 near the national 24-h $\text{PM}_{2.5}$ standard ($35 \mu\text{gm}^{-3}$) (AirKorea). The full co-evolution of PM_1 , gaseous
660 pollutants, and meteorological conditions is provided in Fig. S6, showing an average ambient temperature
661 of $1.76 \pm 4.3^\circ\text{C}$ and average relative humidity (RH) of $56.9 \pm 17.5\%$ during the study.

662 Figure 1 summarizes the overall non-refractory submicron aerosol (NR-PM1) composition and the
663 identified OA factors.”

664 Section 3.1.1: The identification of nitrogen-containing organic aerosols (NOA) could be better supported.
665 I encourage the authors to provide additional evidence, for instance through mass spectral comparison with
666 previous studies, or by applying the NO/NO_2 ratio approach to assess NOA, and then comparing the results
667 with PMF-based identification.

668 Thank you for the helpful suggestions. We agree that the identification of NOA must be clearly supported.
669 Regarding the $\text{NO}^+/\text{NO}_2^+$ ratio: This metric is a well-established diagnostic for assessing the thermal
670 decomposition (and thus functionality) of inorganic or organic nitrate (NO_3^-) in the AMS nitrate channel.
671 It is not applicable to reduced-nitrogen amines ($\text{R}-\text{NH}_2$) that define NOA, which are detected in the organic
672 spectrum as reduced $\text{C}_x\text{H}_y\text{N}^+$ fragments (e.g., m/z 30, 44, 58, 86). Accordingly, we did not apply the
673 $\text{NO}^+/\text{NO}_2^+$ method to the NOA factor. In order to strengthen the mass spectral evidence as requested, we
674 have enhanced the discussion in 3.1.1 to clearly emphasize the spectral matching of our NOA factor against
675 established literature reference spectra of amines (Fig. 3). Now the relevant section reads:

676 “The NOA factor exhibited the highest nitrogen-to-carbon (N:C) ratio (0.22) and the lowest oxygen-to-
677 carbon (O:C) ratio (0.19) among all POA factors (Fig. S2), indicating a chemically reduced, nitrogen-rich
678 composition. The factor represents semi-volatile, reduced nitrogen species that originate from primary
679 urban combustion sources but whose observed mass in the particle phase is enhanced by rapid secondary
680 partitioning and salt formation (Ge et al., 2011; You et al., 2014). The NOA mass spectrum was dominated
681 by amine-related fragments including m/z 30 (CH_4N^+), 44 ($\text{C}_2\text{H}_6\text{N}^+$), 58 ($\text{C}_3\text{H}_8\text{N}^+$), and 86 ($\text{C}_5\text{H}_{12}\text{N}^+$)
682 (Fig. 3a). The spectral signature of the factor is defined by the characteristic dominance of the m/z 44
683 fragment, which typically serves as the primary marker for dimethylamine (DMA)-related species, closely
684 followed by m/z 58 (trimethylamine, TMA) and m/z 30 (methylamine, MA). This profile is in strong
685 agreement with NOA factors resolved via PMF in other polluted environments. For instance, the dominance
686 of m/z 44 and m/z 30 aligns with amine factors reported in New York City (Sun et al., 2011) and Pasadena,
687 California (Hayes et al., 2013). This DMA-dominated signature is also consistent with seasonal
688 characterization of organic nitrogen in Beijing (Xu et al., 2017) and Po Valley, Italy (Saarikoski et al.,
689 2012), reinforcing the common chemical signature of reduced organic nitrogen across diverse urban and
690 regional environments. Furthermore, the presence of non-negligible signals at m/z 58 and m/z 86 supports
691 the contribution of slightly larger alkylamines, a pattern that aligns well with established AMS laboratory
692 reference spectra for these reduced nitrogen compounds (Ge et al., 2011; Silva et al., 2008)”

693 Line 187-188: Please clarify whether the identified NOA is of primary or secondary origin.

694 Thank you for requesting this clarification. The NOA factor exhibits characteristics of both primary
695 emissions and rapid secondary processing, which is typical for reduced nitrogen species. We have clarified
696 in the manuscript that the factor is best characterized as a semi-volatile component derived primarily from
697 combustion emissions whose particle-phase concentration is enhanced by atmospheric processing.

698 We base this conclusion on the following evidence.

- 699 1. Primary Origin (Precursors): The mass spectrum (Fig. 3) and elemental ratios (O:C=0.19, low
700 oxidation) closely match low-molecular-weight alkylamines (e.g., DMA, TMA), which are
701 primarily emitted during high-temperature combustion of nitrogen-rich fuels (Ge et al., 2011). The
702 night-early morning peaks (Fig. 2) further link it to near-source urban combustion/heating activities.
- 703 2. Secondary Enhancement (Observed Mass): The observed high particle-phase concentration is
704 strongly influenced by secondary processes:
 - 705 ○ Partitioning: Amines are semi-volatile and highly basic. Their particle-phase retention
706 (NOA mass) relies on partitioning facilitated by low temperature and reaction with acidic
707 species to form low-volatility aminium salts (a secondary process).
 - 708 ○ Meteorological Dependence: We observed a strong increase in NOA when relative
709 humidity (RH) surpassed 60% (Fig. 2), suggesting that RH-driven partitioning and
710 enhanced formation of aminium salts are critical for its detection (Milic et al., 2016).

711 We have revised the text to clearly articulate this dual nature, emphasizing that while the precursor is
712 primary, the particle mass is semi-secondary (governed by partitioning).

713 "The NOA factor exhibited the highest nitrogen-to-carbon (N:C) ratio (0.22) and the lowest oxygen-to-
714 carbon (O:C) ratio (0.19) among all POA factors (Fig. S2), indicating a chemically reduced, nitrogen-rich
715 composition. **The factor represents semi-volatile, reduced nitrogen species that originate from**
716 **primary urban combustion sources but whose observed mass in the particle phase is enhanced by**
717 **rapid secondary partitioning and salt formation (Ge et al., 2011; You et al., 2014).**

718 Line 212-213: The abbreviations for OOAAs have already been introduced earlier.

719 We removed repeated definitions of OOA and OA.

720 Line 267: The abbreviation "OA" has also been defined earlier.

721 We removed repeated definitions of OOA and OA.

722 Line 268: ... observations at where?

723 Thank you for pointing this out. We had not explicitly stated the location. The observations were made in
724 Seoul, Korea. We have clarified the sentence accordingly.

725 "This trend aligns with previously reported spring and fall observations in Seoul, Korea (Kang et al., 2022;
726 Jeon et al., 2023)."

727 Line 292: Section 3.3?

728 Corrected. Thanks.

729 Line 302-303: Please add a few references to situate your results in the context of previous literature.

730 We added references demonstrating that functionalization typically lowers volatility, while fragmentation
731 increases it and can reduce SOA yields despite higher O:C; this supports our emphasis on pairing
732 composition with direct volatility constraints. Now the section reads:

733 “Generally, the oxygen-to-carbon (O:C) ratio of organic aerosols (OA) is inversely related to their volatility.
734 As O:C increases through aging, the effective saturation concentration (C*) typically decreases, resulting
735 in lower volatility (Donahue et al., 2006; Jimenez et al., 2009). This common relationship arises because
736 the addition of oxygen-containing functional groups (e.g., hydroxyl, carboxyl, carbonyl), which increases
737 molecular weight and enhances intermolecular interactions such as hydrogen bonding, thereby reducing
738 vapor pressure (Jimenez et al., 2009; Kroll and Seinfeld, 2008). Moreover, oxidative aging often leads to
739 oligomerization or functionalization, promoting particle-phase retention and reducing the effective
740 saturation concentration (C*) (Donahue et al., 2011; Robinson et al., 2007). However, in this study, the
741 most oxidized OA factor—MO-OOA, with a high O:C ratio of 1.15—exhibited unexpectedly high
742 volatility. Its volatility distribution was skewed toward SVOCs and IVOCs (Fig. 5), and its rapid mass loss
743 in MFR thermograms (Fig. S9) further indicated low thermal stability. This observation appears to
744 contradict the usual inverse O:C–volatility relationship; however, under winter haze conditions—with
745 suppressed O₃/low OH, particle-phase autoxidation and fragmentation can yield higher-O:C yet more
746 volatile products, with enhanced condensation on abundant particle surface area (details below).”

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