

Supplement of

# Measurement report: High contribution of N<sub>2</sub>O<sub>5</sub> uptake to particulate nitrate formation in NO<sub>2</sub>-limited urban areas

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## Supplementary Texts

### S1. Measurements techniques.

The O<sub>3</sub>, NO<sub>x</sub>, PM<sub>2.5</sub>, SO<sub>2</sub>, and CO were monitored by Thermo Instrument 49i, 17i, 1405DF, 43i, and 48i (Thermo Fisher Scientific, Waltham, MA, USA), respectively. Water-soluble gas (HONO) and water-soluble inorganic ions (sulfate, nitrate, ammonium and chloride) were measured by an online analyzer (MARGA ADI 2080, Applikon Analytical B.V, Netherlands). Organic carbon and elemental carbon in PM<sub>2.5</sub> were measured by OC/EC analyzer (Model 4, Sailhero Environmental Protection High-tech Co., Heibei, China). Around 106 species of VOCs were observed in hourly interval by a gas chromatography-mass spectrometer (GC-FID/MS, TH-300B, Wuhan, CN). As for HCHO and PAN, they were measured by the HCHO analyzer (FMS-100, Focused Photonics Inc., Hangzhou, China) and the PAN analyzer (PANs-1000, Focused Photonics Inc., Hangzhou, China), respectively. In addition, ambient temperature (T), relative humidity (RH), atmospheric pressure (P), wind speed (WS), and wind direction (WD) were measured by an integrated sensor (150WX, Airmar, Milford, NH, USA). The photolysis frequencies (JO<sub>1</sub>D, JNO<sub>2</sub>, JHONO, JNO<sub>3</sub>, JHCHO and JH<sub>2</sub>O<sub>2</sub>) were measured by a photolysis spectrometer (FPS-100, Focused Photonics Inc., Hangzhou, China). The uncertainty and detection limit of above observing instruments were described in our previous studies (Liu et al., 2022a; Liu et al., 2023; Hu et al., 2022; Hu et al., 2020; Liu et al., 2022b; Yang et al., 2023).

### S2. Detailed settings of Iodide-TOF-CIMS and the calibrations of N<sub>2</sub>O<sub>5</sub> and ClNO<sub>2</sub>.

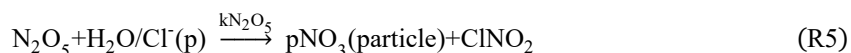
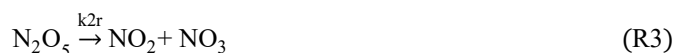
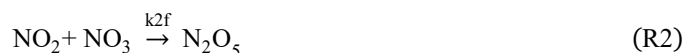
In this work, the chemical ionization atmospheric pressure interface long-time-of flight mass spectrometer (API-ToF-CIMS, Aerodyne Research Inc., USA) with an iodide source was used to detect ambient N<sub>2</sub>O<sub>5</sub> and ClNO<sub>2</sub>. A stainless-steel tube with a length of near 2 meters and a diameter of 1/4 inch inner diameter was used for sampling. The sampling flow rate was set as 10 standard liters per minute (SLPM) and only 2SLPM from the total flow was diverted to the CIMS. The nitrogen (N<sub>2</sub>) flow (99.999%, 2.7 SLPM) containing the methyl iodide gas (CH<sub>3</sub>I) released by a heated CH<sub>3</sub>I permeation tube cylinder, passed through a soft X-ray device (Tofwerk AG, P-type) to generate reagent ions I<sup>+</sup>. The I<sup>+</sup> was combined with the target gas in ion molecule reaction (IMR) chamber and then detected by ToF-CIMS instrument. The ambient N<sub>2</sub>O<sub>5</sub> and ClNO<sub>2</sub> were detected as I(N<sub>2</sub>O<sub>5</sub>)<sup>+</sup> and I(ClNO<sub>2</sub>)<sup>+</sup> clusters at 235 and 238 m/z. In order to minimize the effect of particles deposited on the surface of the sampling inlet, the tube was cleaned by deionized water and dried by nitrogen flow once a week.

The calibration of  $\text{N}_2\text{O}_5$  and  $\text{ClNO}_2$  was set similar to previous studies (Wang et al., 2022b; Wang et al., 2022a; Thaler et al., 2011). And the calibration of  $\text{N}_2\text{O}_5$  is based on the already calibrated  $\text{ClNO}_2$ . Briefly, during the  $\text{ClNO}_2$  calibration process, a  $\text{N}_2$  flow ( $50 \text{ mL min}^{-1}$ ) containing 6 ppmv  $\text{Cl}_2$  was passed over a slurry containing  $\text{NaNO}_2$  and  $\text{NaCl}$  to produce  $\text{ClNO}_2$  and then the mixed flow was sampled into the CIMS instrument under a certain RH. The function of  $\text{NaCl}$  was minimizing formation of  $\text{NO}_2$  as a byproduct. To quantify  $\text{ClNO}_2$ , the mixed flow was fed directly into a cavity attenuated phase shift spectroscopy instrument (CAPS) or passed through a thermal dissociation tube at  $380^\circ\text{C}$  before being fed into the CAPS to measure the background  $\text{NO}_2$  and the total  $\text{NO}_2$  (background  $\text{NO}_2$  and the  $\text{NO}_2$  decomposed by  $\text{ClNO}_2$ ), respectively. The difference between the total  $\text{NO}_2$  and background  $\text{NO}_2$  was equivalent to the concentration of  $\text{ClNO}_2$ . And the normalized signal of  $\text{ClNO}_2$  was tested in different RH conditions. To calibrate the  $\text{N}_2\text{O}_5$ , the stable flow of  $\text{N}_2\text{O}_5$  was produced by mixing the  $\text{O}_3$  flow and  $\text{NO}_2$  flow. The  $\text{O}_3$  oxidized  $\text{NO}_2$  to form  $\text{NO}_3$ , which was further oxidized to form  $\text{N}_2\text{O}_5$ . Under different certain RH, the stable- $\text{N}_2\text{O}_5$  flow was sampled into CIMS instrument to obtain a normalized signal for  $\text{N}_2\text{O}_5$  varying with RH. Subsequently, we delivered the  $\text{N}_2\text{O}_5$  flow at  $\sim 50\%$  RH through a supersaturated sodium chloride solution to convert  $\text{N}_2\text{O}_5$  to  $\text{ClNO}_2$  and then sampled in to CIMS. The concentration of  $\text{N}_2\text{O}_5$  was delivered at  $\sim 50\%$  RH by already calibrated  $\text{ClNO}_2$  and then scaled to other RH conditions by the normalized  $\text{N}_2\text{O}_5$  signal, which was mentioned above. The final calibration files for  $\text{N}_2\text{O}_5$  and  $\text{ClNO}_2$  at different RH are shown as **Figure S2**. The detection limit of  $\text{N}_2\text{O}_5$  and  $\text{ClNO}_2$  were 1.3 and 0.43 ppt, respectively.

### S3. The interactive box model.

The interactive box model considers five main chemical reactions related to  $\text{N}_2\text{O}_5$  (see **R1 – R5**). Reaction rate of  $k_1$  and reversible reaction  $k_{\text{eq}}$  are  $1.2 \times 10^{-13} e^{-2450/T}$  and  $2.7 \times 10^{-27} e^{11000/T}$ , respectively. (Demore et al., 1997; Brown et al., 2003b) Measured  $\text{O}_3$ ,  $\text{NO}_2$ ,  $\text{N}_2\text{O}_5$ ,  $\text{ClNO}_2$ , VOCs and meteorological parameters during the observation period were used to constrain the model simulation. The hourly  $\text{NO}_3$  reactivity with VOCs was estimated from the sum of measured VOCs concentrations and reaction rate coefficients of MCM v3.3.1 (Atkinson and Arey, 2003; Wang et al., 2017; Jenkin et al., 2003; Jenkin et al., 2015). Due to the lack of measurements for biogenic monoterpenes, the calculated  $\text{NO}_3$  reactivity could be considered as a lower limit. Procedure of this model, as presented in other studies (Chen et al., 2020; Niu et al., 2022), includes a backward integration step to get simulated concentration

of O<sub>3</sub> and NO<sub>2</sub>, a forward integration step to get simulated concentration of NO<sub>3</sub> and N<sub>2</sub>O<sub>5</sub>, a calculating the first-order loss rate of N<sub>2</sub>O<sub>5</sub> (kN<sub>2</sub>O<sub>5</sub>) step and an adjusting step to reduce the difference between simulated and measured N<sub>2</sub>O<sub>5</sub> to less than 1%. The ClNO<sub>2</sub> production rate (φClNO<sub>2</sub>) was calculated by **eq 1**. The initial time for each model simulation was set to the sunset time, with a total duration of 24 hours.



#### S4. The steady-state method.

The input data was also from the 2022 winter observation. Based on the basic framework of the nighttime N<sub>2</sub>O<sub>5</sub> chemistry described in **R1 – R5** and 5-minute resolution data of N<sub>2</sub>O<sub>5</sub>, NO<sub>2</sub>, O<sub>3</sub>, and T, the lifetime of N<sub>2</sub>O<sub>5</sub> (τN<sub>2</sub>O<sub>5</sub>) is estimated by the **eq 3** with K<sub>eq</sub> (calculated by the **eq 2**) in assumption of the steady state (Brown et al., 2003a). The robust kN<sub>2</sub>O<sub>5</sub> could be derived by the **eq 4** as an intercept (Chen et al., 2022). Similarly, the (φClNO<sub>2</sub>) was calculated by the **eq 1**. Notably, the results are summarized in **Table S6**.

$$\phi\text{ClNO}_2 = \frac{[\text{ClNO}_2]}{\int k_{\text{N}_2\text{O}_5} [\text{N}_2\text{O}_5] dt} \quad (1)$$

$$K_{\text{eq}}(T) = 2.7 \times 10^{-27} e^{11000/T} \quad (2)$$

$$\tau_{\text{N}_2\text{O}_5} = \frac{[\text{N}_2\text{O}_5]}{K_1(T)[\text{NO}_2][\text{O}_3]} \quad (3)$$

$$(\tau_{\text{N}_2\text{O}_5})^{-1} = k_{\text{N}_2\text{O}_5} + \frac{k_{\text{NO}_3}}{K_{\text{eq}}[\text{NO}_2]} \quad (4)$$

#### S5. Setting of the multiphase chemical box model

The F0AM model was used to simulated the formation rate of pNO<sub>3</sub><sup>-</sup> by different formation pathways. Measured hourly interval data of trace gases (VOCs, O<sub>3</sub>, NO, NO<sub>2</sub>, CO, SO<sub>2</sub>), photochemically active species (HONO, HCHO, PAN), meteorological variables (temperature, relative humidity,

atmospheric pressure, and photolysis frequencies), reanalysis data (boundary layer height, surface albedo and overhead ozone column) and simulated heterogeneous parameters in wintertime of 2022 as summarized in **Table S5** were used to constrain this model. It should be noted that the photolysis rate constants of the model parameterization were corrected by the observed and simulated photolysis rates of NO<sub>2</sub>, except observed photolysis frequencies. To better simulate the concentration of intermediate species, 3-day spin-up was set before each model simulation. In addition to chemical processes, physical loss processes including dry deposition and dilution were also considered in this model including dry deposition and dilution process. The dry deposition velocity for the constrained species was set as Liu et al (Liu et al., 2022b). The basic dilution rate was  $2 \times 10^{-5} \text{ s}^{-1}$  and varied with the boundary height. The incorporated mechanisms in the F0AM model are presented as **Table S2**.

## **S6. Detailed information about explainable machine learning techniques**

### **(1) Setting of XGBoost model and validation.**

Extreme gradient boosting (XGBoost) model was used to explore the influencing factors of N<sub>2</sub>O<sub>5</sub> uptake. The input data for the model were hourly data at night (18:00 – 06:00 the next day) during winter 2022. The dataset including 813 samples was divided into a training set (80%) and a test set (20%). The hyperparameters were tuned using grid search and cross-validation methods. Specifically, for a single hyperparameter, grid search was used to obtain its more appropriate value range. For combinations of hyperparameters, the whole training set was split into ten folds. Grid search was then conducted over the pre-adjusted combinations of hyperparameters by training on nine folds and predicting on the tenth fold, following a cross-validation procedure. The hyperparameters of the final model were determined according to the R<sup>2</sup> function score. For the key hyperparameters of the XGBoost model in this study, the number of trees was 100, the learning rate was 0.07, the max depth was 7 and the regularization parameter lambda was set as 1.38.

The established model was evaluated by three statistical indicators including R<sup>2</sup> value, Mean Absolute Error (MAE) and Root Mean Square Error (RMSE). The R<sup>2</sup> is used to judge the goodness of fit between the predicted and observed results. The MAE provides a measure of predicted and observed results. And the RMSE characterizes the predictive accuracy of the model. The equations to calculate R<sup>2</sup>, MAE and RMSE are **eq 5**, **eq 6**, and **eq 7**, respectively. The XGBoost model built in this study performed well, with an R<sup>2</sup> of 0.97, an MAE of 0.07, and an RMSE of 0.14 (**Figure S5**).

$$R^2 = 1 - \frac{\sum_{i=1}^N (y_i - y'_i)^2}{\sum_{i=1}^N (y_i - \bar{y})^2} \quad (5)$$

$$MAE = \sqrt{\frac{\sum_{i=1}^N (y_i - y'_i)^2}{N}} \quad (6)$$

$$RMSE = \frac{1}{N} \sum_{i=1}^N |y_i - y'_i|^2 \quad (7)$$

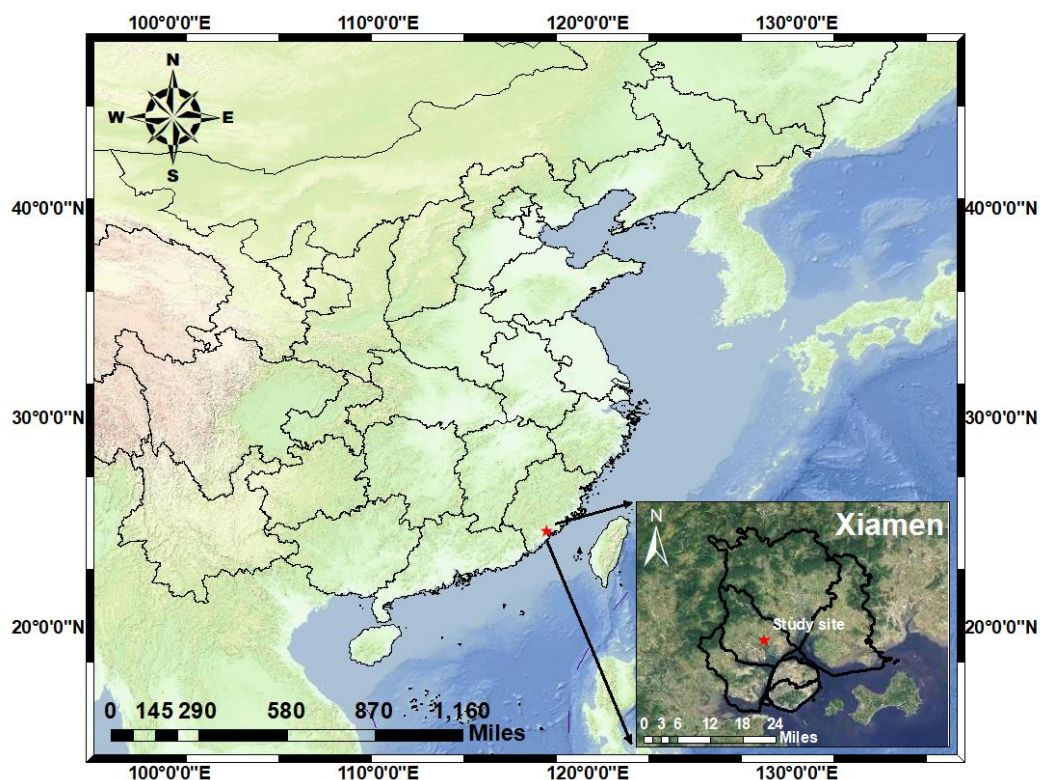
Where  $y_i$ ,  $y'_i$ ,  $\bar{y}$ , and  $N$  refer to the actual value of the  $i$ -th sample, the simulated value of the  $i$ -th sample, the mean of the actual values of all samples, and  $N$  is the number of samples, respectively.

## (2) The Shapley Additive Explanation theory

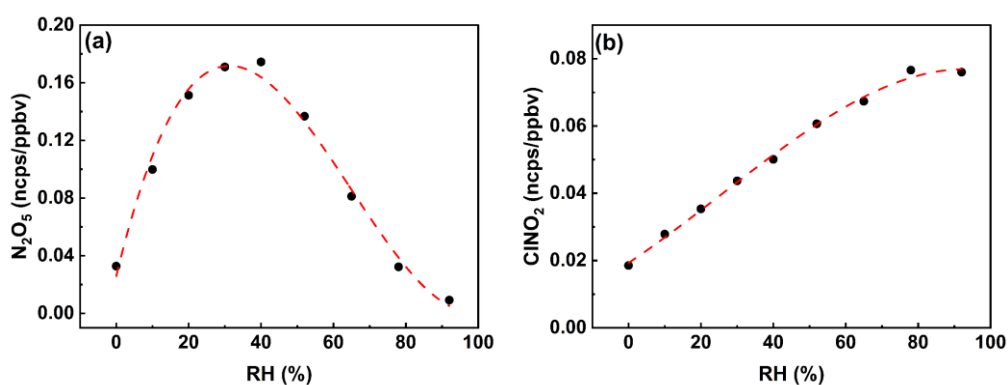
In this study, the target of SHAP method is to find the drivers of  $N_2O_5$  uptake. The SHAP method could quantified the contribution of each feature to the variation of  $N_2O_5$  uptake. Based on the coalitional game theory (Shapley, 1953), SHAP values were computed according to eq 8 and describe the contribution of each feature to the prediction of an individual sample relative to the base prediction value of all samples.

$$g(x_i) = \phi(x)_{base} + \sum_{j=1}^M \phi(x_{i,j}) \quad (8)$$

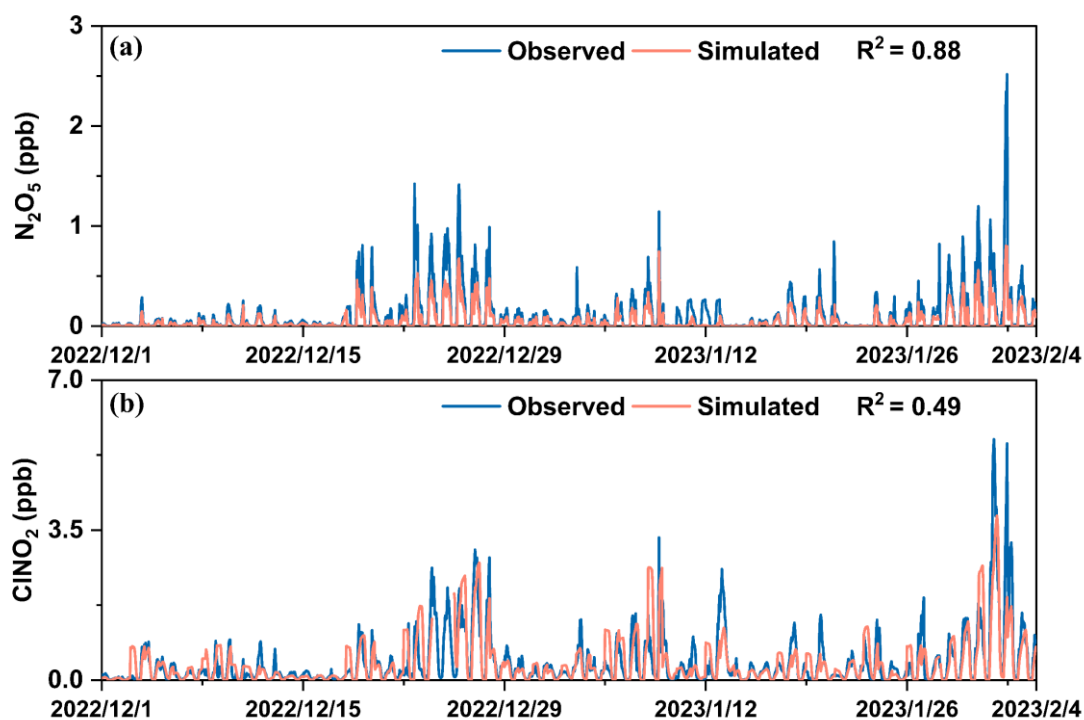
Where  $g(x_i)$  is a predicted value for each sample  $x_i$  containing  $M$  features,  $\phi(x)_{base}$ , the base value, represents the predicted value for all samples.  $\phi(x_{i,j})$  is the SHAP value of the feature  $j$  in sample  $x_i$  showing its impact on the prediction. In this study, the importance of input variables to variation of  $N_2O_5$  uptake was quantified by absolute SHAP values.  $SHAP > 0$  means that the feature has a positive impact to  $N_2O_5$  uptake relative to the base value. While conversely, it exerts a negative contribution.



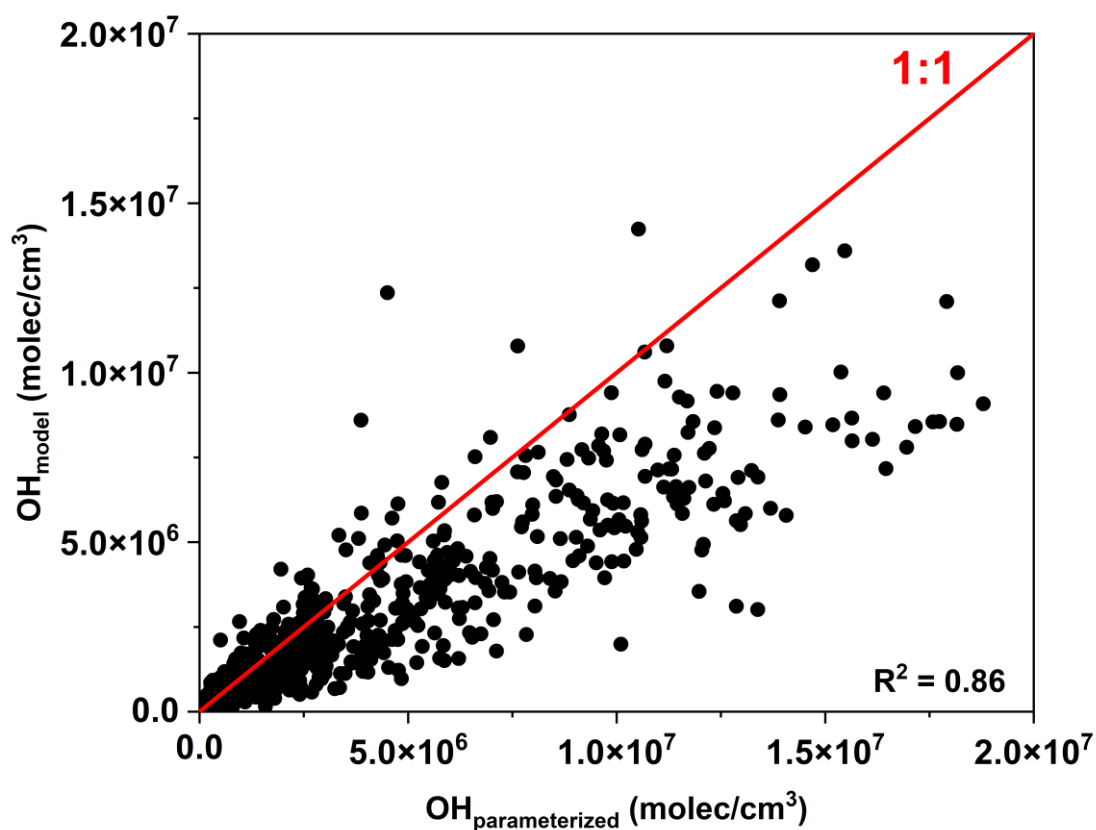
**Figure S1.** Measurement site located in the Atmospheric Observation Supersite of Institute of Urban Environment, Chinese Academy of Science (red star; 24.61°N, 118.06°E) in Xiamen city of Southeast China. The observation site is approximately 70 meters above ground level and surrounded by the educational institutions, residential buildings, and traffic arteries, thereby, representing a typical urban site.



**Figure S2.** The calibration curve of CIMS instrument as a function of relative humidity (RH) for (a)  $N_2O_5$  and (b)  $ClNO_2$ .



**Figure S3.** Time series of the observed and simulated (a)  $\text{N}_2\text{O}_5$  and (b)  $\text{ClNO}_2$  from the multiphase chemical box model, respectively.

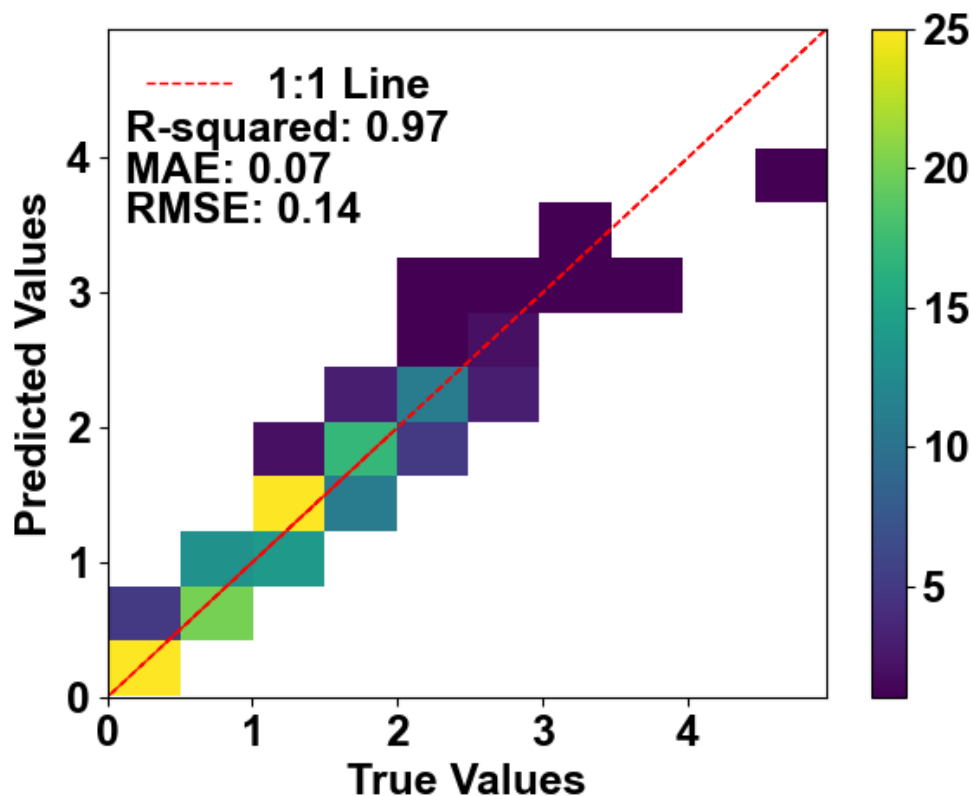


**Figure. S4.** Concentration of OH radical derived from parameterized method and F0AM model

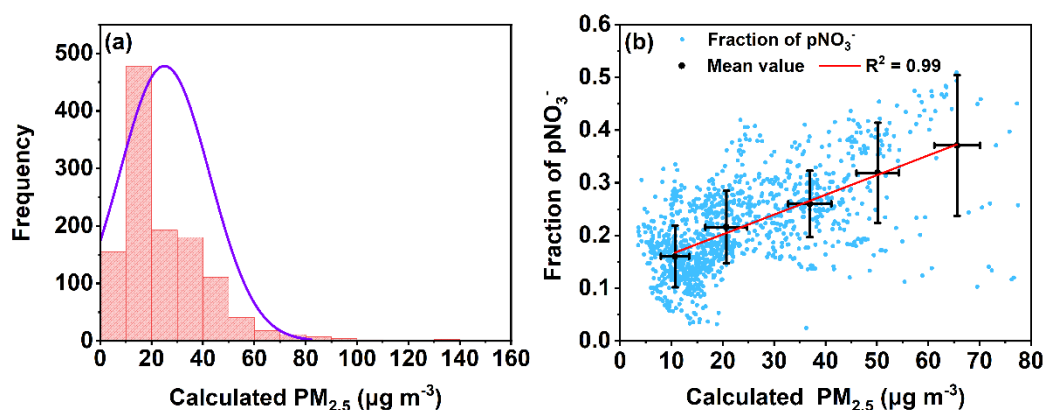


simulation. The method of OH parameterization was suggested by Ehhalt and Rohrer (Ehhalt and Rohrer,

$$2000): [\text{OH}] = 4.1 \times 10^9 \times (\text{J}_{\text{O1D}})^{0.83} \times (\text{J}_{\text{NO2}})^{0.19} \times \frac{140 \times [\text{NO}_2] + 1}{0.41 \times [\text{NO}_2]^2 + 1.7 \times [\text{NO}_2] + 1}$$

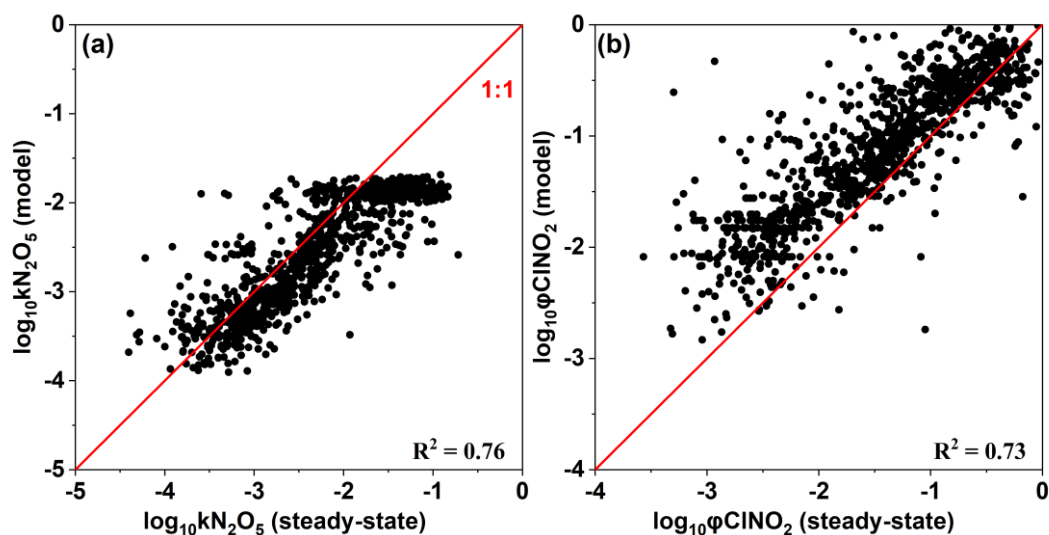


**Figure S5. Performance of the XGBoost model based on the test set.** The color bar is colored with number density.

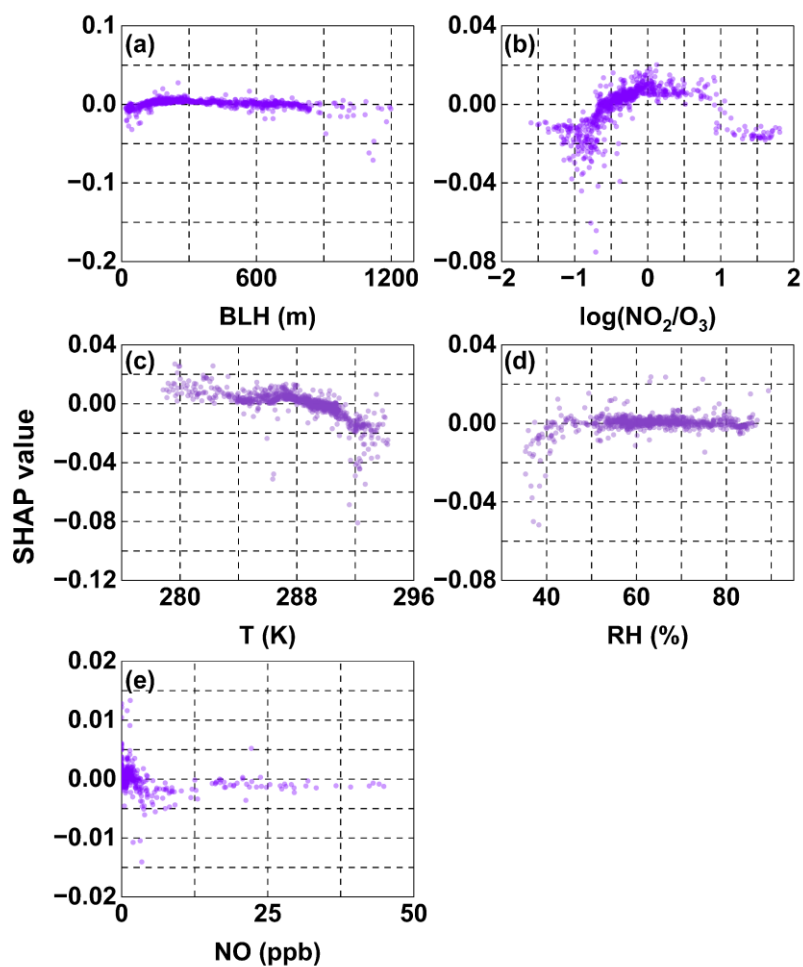


**Figure S6. Histogram of the distribution of hourly calculated  $\text{PM}_{2.5}$  concentrations (a) and variation of nitrate fractions with the levels of the hourly calculated  $\text{PM}_{2.5}$  concentrations (b).** The calculated  $\text{PM}_{2.5}$  concentration is the summed concentration of major chemical components, including water-soluble

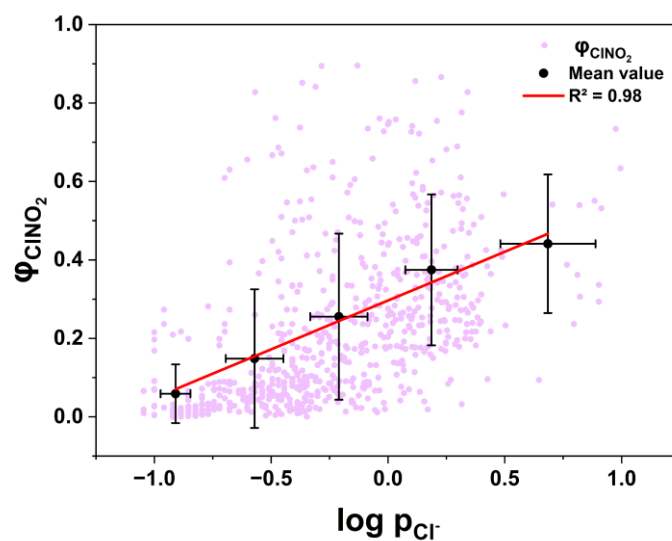
inorganic ions, trace elements, organic carbon and elemental carbon. Error bar represents one standard deviation.



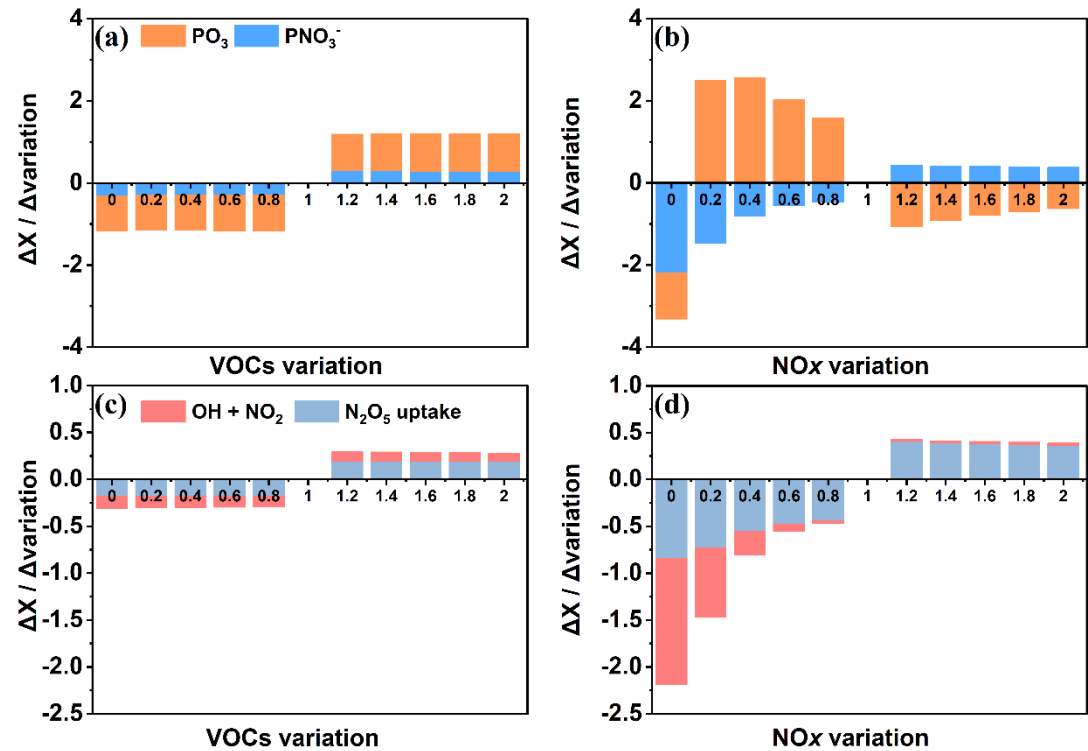
**Figure S7.** Comparison of parameters derived from iterative box model and steady state approximation method. **(a)**  $\text{N}_2\text{O}_5$  heterogeneous loss rate and **(b)**  $\text{ClNO}_2$  production rate, respectively.



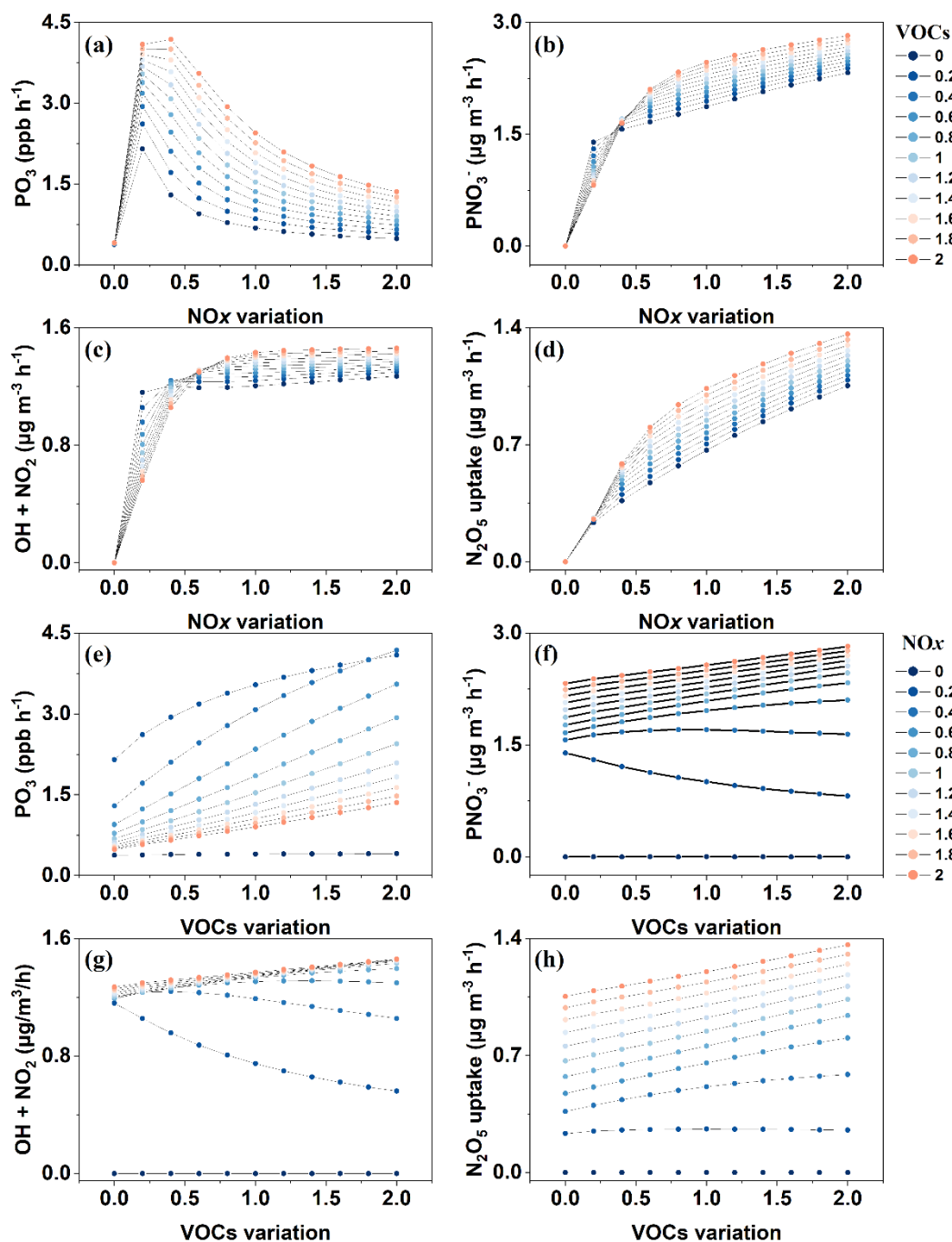
**Figure S8.** Main effect of the factors (a) BLH, (b)  $\log \frac{\text{NO}_2}{\text{O}_3}$ , (c) Temperature, (d) RH, and (e) NO on the  $\text{N}_2\text{O}_5$  uptake obtained by XGBoost-SHAP explainer method.



**Figure S9.** The relationship between the log of particulate chlorine ( $\log(p_{Cl^-})$ ) and  $\phi ClNO_2$  in winter 2022. The error bars represent standard deviation, and the black circles are the mean value of each bin, which are (minimum, -0.8], (0.8, -0.4], (-0.4, 0], (0, 0.4], (0.4, maximum] for  $\log p_{Cl^-}$ .



**Figure S10.** The response strength (RS) of  $PO_3$  and  $PNO_3^-$  (a, b), and the RS of the  $PNO_3^-$  via  $N_2O_5$  uptake and  $OH + NO_2$  (c, d) to different VOCs and NOx change scenarios. It's noted that when VOCs change, the emission rate of NOx is fixed as base, and when NOx change, the emission rate of VOCs is fixed as base.



**Figure S11.** Relationships of mean formation rates of PO<sub>3</sub> (a, e), PNO<sub>3</sub><sup>-</sup> (b, f), N<sub>2</sub>O<sub>5</sub> uptake (c, g), and OH + NO<sub>2</sub> (d, h) to the normalized NOx and VOCs in all the examined scenarios simulated by the multiphase chemical box model. The normalized base was the average case of winter 2022.

**Table S1.** Comparison of the winter NO<sub>2</sub>, O<sub>3</sub> and total VOCs levels in urban areas of China.

Location	NO <sub>2</sub> (ppb)	O <sub>3</sub> (ppb)	VOCs (ppb)	PM <sub>2.5</sub> (μg/m <sup>3</sup> )	Observation period	Reference
Xiamen	10.9	27.3	18.2	14.3	Dec,2022 – Feb, 2023	This study

Beijing	-	-	35.8	-	Jan, 2019	(Wei et al., 2021)
Beijing	19.1	9.2	41.2	55.2	Dec, 2018 – Jan, 2019	(Liu et al., 2021b)
Guangzhou	22.9	30.7	33.1	35.5	Jan, 2021	(Cheng et al., 2024)
Nanjing	-	-	40.2	-	Winter, 2016	(Zhao et al., 2020)
Shanghai	22.6	14.1	63.6	34.0	Dec, 2017 – Jan, 2018	(Liu et al., 2021a)
Chengdu	-	-	53.3	-	Jan, 2019	(Xiong et al., 2021)

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228 **Table S2.** The incorporated mechanism in the F0AM model.

Reaction	Rate constant and emission factor	Reference
$\text{HNO}_3 \rightarrow \text{pNO}_3^-$	gas-particle partition coefficient	(Zare et al., 2018)
$\text{N}_2\text{O}_5 + \text{aerosol} + \text{H}_2\text{O} \rightarrow \text{pNO}_3^-$	$k\text{N}_2\text{O}_5 \times (2 - \phi\text{ClNO}_2)$	(Chen et al., 2020)
$\text{N}_2\text{O}_5 + \text{aerosol} + \text{pCl}^- \rightarrow \text{ClNO}_2$	$\phi\text{ClNO}_2$	(Chen et al., 2020)
$\text{ClNO}_2 + \text{h}\nu \rightarrow \text{Cl} + \text{NO}_2$	Based on IUPAC	(Sherwen et al., 2016)
Other halogen mechanisms as GEOS-Chem	Based on IUPAC	(Sherwen et al., 2016)

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230 **Table S3.** Observed period of different winters from 2019 to 2023.

Year	Sampling period
2019	2020.01.01–2020.01.26
2020	2020.12.01–2020.12.14
2021	2022.01.0–2022.02.05
2022	2022.12.01–2023.02.04
2023	2023.12.01–2023.12.31

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232 **Supplementary Table 4.** Summary of the key parameters of nighttime  $\text{NO}_3\text{-N}_2\text{O}_5$  chemistry around the  
233 world. Values of parameters are mean level unless stated.

Region	Site type	Study period	$k\text{N}_2\text{O}_5 \text{ het (s}^{-1}\text{)}$	$\phi\text{ClNO}_2$	Reference
Xiamen, China	Urban	Dec,2022 – Feb,2023	$7.64 \times 10^{-3}$	0.193	This study
Beijing, China	Rural	May, 2016 – Jun, 2016	$8.1 \times 10^{-4}$	$0.73 \pm 0.25$	(Wang et al., 2018)
Beijing, China	Urban (vertical)	Oct, 2020 – Nov, 2020	$1.42 \times 10^{-3}$	-	(Ma et al., 2023)

Hebei, China	Semi-rural	Jun, 2014 – Jul, 2014	$1.3 \times 10^{-2}$	0.30	(Tham et al., 2016)
Guangdong, China	Semi-rural	Jan, 2017	$3.78 \times 10^{-3}$ – $9.00 \times 10^{-3}$	0.18 – 0.32	(Yun et al., 2018)
UK	Flight track	Jul, 2010	$9.3 \times 10^{-5}$ – 0.001	-	(Morgan et al., 2015)

**Table S5. Mean concentrations and standard deviations of chemical species and meteorological parameters during the observation period.** Concentrations are in ppbv unless stated. The daytime is 07:00 – 17:00 (local time).

Species	Mean	Standard deviation
N <sub>2</sub> O <sub>5</sub>	0.110	0.215
ClNO <sub>2</sub>	0.421	0.644
O <sub>3</sub>	27.3	14.1
NO	2.83	6.59
NO <sub>2</sub>	10.9	6.43
CO	0.409	0.101
SO <sub>2</sub>	1.95	0.537
HCHO	1.69	0.858
HONO	0.568	0.519
PAN	0.409	0.302
T (K)	290	4.07
RH (%)	58.3	14
BLH (m)	513	339
P (Pa)	1000	76.6
Calculated PM <sub>2.5</sub> (µg/cm <sup>3</sup> )	14.3	9.43
C <sub>2</sub> H <sub>6</sub>	3.04	0.731
C <sub>3</sub> H <sub>8</sub>	2.74	1.43
NC <sub>4</sub> H <sub>10</sub>	1.19	0.845
IC <sub>4</sub> H <sub>10</sub>	0.808	0.531
NC <sub>5</sub> H <sub>12</sub>	0.302	0.224
IC <sub>5</sub> H <sub>12</sub>	0.673	0.75
C <sub>2</sub> H <sub>4</sub>	1.14	0.689
C <sub>3</sub> H <sub>6</sub>	0.313	0.323
C <sub>4</sub> H <sub>6</sub>	0.0137	0.0174
CBUT <sub>2</sub> ENE	0.0184	0.0324
TBUT <sub>2</sub> ENE	0.0139	0.0376
BUT1ENE	0.0641	0.0791
PENT1ENE	0.00471	0.0119
C <sub>3</sub> H <sub>8</sub>	0.00724	0.0142
CHEX	0.0205	0.0199
NC <sub>6</sub> H <sub>14</sub>	0.107	0.154

M <sub>3</sub> PE	0.0751	0.0871
M <sub>2</sub> PE	0.0674	0.0827
M <sub>23</sub> C <sub>4</sub>	0.0697	0.108
M <sub>22</sub> C <sub>4</sub>	0.0261	0.0262
NC <sub>7</sub> H <sub>16</sub>	0.0472	0.0765
M <sub>3</sub> HEX	0.0488	0.0888
M <sub>2</sub> HEX	0.0334	0.0568
NC <sub>8</sub> H <sub>18</sub>	0.0131	0.0148
NC <sub>9</sub> H <sub>20</sub>	0.00833	0.00986
NC <sub>10</sub> H <sub>22</sub>	0.00589	0.00652
NC <sub>11</sub> H <sub>24</sub>	0.00588	0.00742
NC <sub>12</sub> H <sub>26</sub>	0.02	0.145
HEX1ENE	0.00208	0.00474
C <sub>2</sub> H <sub>2</sub>	1.09	0.406
BENZENE	0.251	0.107
TOLUENE	0.817	0.992
EBENZ	0.148	0.183
OXYL	0.0888	0.126
PXYL	0.328	0.475
STYRENE	0.0284	0.0663
PBENZ	0.00438	0.00352
IPBENZ	0.00285	0.00374
TM <sub>135</sub> B	0.00529	0.00656
TM <sub>124</sub> B	0.0177	0.0231
TM <sub>123</sub> B	0.005	0.00596
OETHTOL	0.00583	0.00602
METHTOL	0.0105	0.0108
PETHTOL	0.00606	0.0063
ACR	0.0583	0.0417
CH <sub>3</sub> COCH <sub>3</sub>	1.38	0.743
MEK	0.339	0.321
IPROPOL	0.106	0.152
ETHACET	0.873	1.15
MTBE	0.0939	0.142
CH <sub>3</sub> CL	0.0125	0.0132
CH <sub>2</sub> CL <sub>2</sub>	1.26	1.77
CHCL <sub>3</sub>	0.038	0.0272
CH <sub>2</sub> CLCH <sub>2</sub> CL	0.279	0.202
CH <sub>3</sub> CH <sub>2</sub> CL	0.0156	0.00755
CL <sub>12</sub> PROP	0.101	0.166
TCE	0.0231	0.0246
TRICLETH	0.0111	0.0104
$j_{NO_2}(s^{-1})$ (daytime)	$3.76 \times 10^{-3}$	$3.02 \times 10^{-3}$
$j_{O_1D}(s^{-1})$ (daytime)	$1.28 \times 10^{-5}$	$1.25 \times 10^{-5}$



$j_{HONO}(s^{-1})$ (daytime)	$6.66 \times 10^{-4}$	$5.33 \times 10^{-4}$
$j_{HCHO}(s^{-1})$ (daytime)	$1.79 \times 10^{-5}$	$1.47 \times 10^{-5}$

**Table S6. The average values of key parameters obtained from the interactive box model and the multiphase chemical box model.** The daytime in this study is from 7:00 to 17:00.

Parameters	Mean $\pm$ standard error
$kNO_3(s^{-1})$	$0.00288 \pm 0.00585$
$kN_2O_5(s^{-1})$	$0.00764 \pm 0.00612$
$\frac{kNO_3}{K_{eq}[NO_2]}(s^{-1})$	$0.000136 \pm 0.000306$
$\phi ClNO_2$	$0.193 \pm 0.284$
OH radicals (daytime) (molec $m^{-3}$ )	$2.52 \times 10^6 \pm 2.84 \times 10^6$
HO <sub>2</sub> radicals (daytime) (molec $m^{-3}$ )	$4.24 \times 10^7 \pm 5.20 \times 10^7$

## References:

- Atkinson, R. and Arey, J.: Atmospheric degradation of volatile organic compounds, Chemical Reviews. 103, 4605-4638, <https://doi.org/10.1021/cr0206420>, 2003.
- Brown, S. S., Stark, H., and Ravishankara, A. R.: Applicability of the steady state approximation to the interpretation of atmospheric observations of NO<sub>3</sub> and N<sub>2</sub>O<sub>5</sub>: art. no. 4539, Journal of Geophysical Research-Atmospheres. 108, <https://doi.org/10.1029/2003jd003407>, 2003a.
- Brown, S. S., Stark, H., Ryerson, T. B., Williams, E. J., Nicks, D. K., Trainer, M., Fehsenfeld, F. C., and Ravishankara, A. R.: Nitrogen oxides in the nocturnal boundary layer:: Simultaneous in situ measurements of NO<sub>3</sub>, N<sub>2</sub>O<sub>5</sub>, NO<sub>2</sub>, NO, and O<sub>3</sub> -: art. no. 4299, Journal of Geophysical Research-Atmospheres. 108, <https://doi.org/10.1029/2002jd002917>, 2003b.
- Chen, X. R., Wang, H. C., and Lu, K. D.: Interpretation of NO<sub>3</sub>-N<sub>2</sub>O<sub>5</sub> observation via steady state in high-aerosol air mass: the impact of equilibrium coefficient in ambient conditions, Atmospheric Chemistry and Physics. 22, 3525-3533, <https://doi.org/10.5194/acp-22-3525-2022>, 2022.
- Chen, X. R., Wang, H. C., Lu, K. D., Li, C. M., Zhai, T. Y., Tan, Z. F., Ma, X. F., Yang, X. P., Liu, Y. H., Chen, S. Y., Dong, H. B., Li, X., Wu, Z. J., Hu, M., Zeng, L. M., and Zhang, Y. H.: Field Determination of Nitrate Formation Pathway in Winter Beijing, Environmental Science & Technology. 54, 9243-9253, <https://doi.org/10.1021/acs.est.0c00972>, 2020.
- Cheng, C. L., Yang, S. X., Yuan, B., Pei, C. L., Zhou, Z. H., Mao, L. Y., Liu, S. L., Chen, D. Y., Cheng, X. Y., Li, M., Shao, M., and Zhou, Z.: The significant contribution of nitrate to a severe haze event in the winter of Guangzhou, China, Science of the Total Environment. 909, <https://doi.org/10.1016/j.scitotenv.2023.168582>, 2024.
- DeMore, W., Howard, C., Sander, S., Ravishankara, A., Golden, D., Kolb, C., Hampson, R., Molina, M., and Kurylo, M.: Chemical Kinetics and Photochemical Data for Use in Stratospheric Modeling, 1997.
- Ehhalt, D. H. and Rohrer, F.: Dependence of the OH concentration on solar UV, Journal of Geophysical Research-Atmospheres. 105, 3565-3571, <https://doi.org/10.1029/1999jd901070>, 2000.
- Hu, B. Y., Liu, T. T., Hong, Y. W., Xu, L. L., Li, M. R., Wu, X., Wang, H., Chen, J. H., and Chen, J. S.: Characteristics of peroxyacetyl nitrate (PAN) in a coastal city of southeastern China: Photochemical mechanism and pollution process, Science of the Total Environment. 719,

<https://doi.org/10.1016/j.scitotenv.2020.137493>, 2020.

Hu, B. Y., Duan, J., Hong, Y. W., Xu, L. L., Li, M. R., Bian, Y. H., Qin, M., Fang, W., Xie, P. H., and Chen, J. S.: Exploration of the atmospheric chemistry of nitrous acid in a coastal city of southeastern China: results from measurements across four seasons, *Atmospheric Chemistry and Physics*. 22, 371-393, <https://doi.org/10.5194/acp-22-371-2022>, 2022.

Jenkin, M. E., Young, J. C., and Rickard, A. R.: The MCM v3.3.1 degradation scheme for isoprene, *Atmospheric Chemistry and Physics*. 15, 11433-11459, <https://doi.org/10.5194/acp-15-11433-2015>, 2015.

Jenkin, M. E., Saunders, S. M., Wagner, V., and Pilling, M. J.: Protocol for the development of the Master Chemical Mechanism, MCM v3 (Part B): tropospheric degradation of aromatic volatile organic compounds, *Atmospheric Chemistry and Physics*. 3, 181-193, <https://doi.org/10.5194/acp-3-181-2003>, 2003.

Liu, T. T., Lin, Y. L., Chen, J. S., Chen, G. J., Yang, C., Xu, L. L., Li, M. R., Fan, X. L., Zhang, F. W., and Hong, Y. W.: Pollution mechanisms and photochemical effects of atmospheric HCHO in a coastal city of southeast China, *Science of the Total Environment*. 859, <https://doi.org/10.1016/j.scitotenv.2022.160210>, 2023.

Liu, T. T., Chen, G. J., Chen, J. S., Xu, L. L., Li, M. R., Hong, Y. W., Chen, Y. T., Ji, X. T., Yang, C., Chen, Y. P., Huang, W. G., Huang, Q. J., and Wang, H.: Seasonal characteristics of atmospheric peroxyacetyl nitrate (PAN) in a coastal city of Southeast China: Explanatory factors and photochemical effects, *Atmospheric Chemistry and Physics*. 22, 4339-4353, <https://doi.org/10.5194/acp-22-4339-2022>, 2022a.

Liu, T. T., Hong, Y. W., Li, M. R., Xu, L. L., Chen, J. S., Bian, Y. H., Yang, C., Dan, Y. B., Zhang, Y. N., Xue, L. K., Zhao, M., Huang, Z., and Wang, H.: Atmospheric oxidation capacity and ozone pollution mechanism in a coastal city of southeastern China: analysis of a typical photochemical episode by an observation-based model, *Atmospheric Chemistry and Physics*. 22, 2173-2190, <https://doi.org/10.5194/acp-22-2173-2022>, 2022b.

Liu, Y. H., Wang, H. L., Jing, S. G., Peng, Y. R., Gao, Y. Q., Yan, R. S., Wang, Q., Lou, S. R., Cheng, T. T., and Huang, C.: Strong regional transport of volatile organic compounds (VOCs) during wintertime in Shanghai megacity of China, *Atmospheric Environment*. 244, <https://doi.org/10.1016/j.atmosenv.2020.117940>, 2021a.

Liu, Z. R., Wang, Y. S., Hu, B., Lu, K. D., Tang, G. Q., Ji, D. S., Yang, X. P., Gao, W. K., Xie, Y. Z., Liu, J. Y., Yao, D., Yang, Y., and Zhang, Y. H.: Elucidating the quantitative characterization of atmospheric oxidation capacity in Beijing, China, *Science of the Total Environment*. 771, <https://doi.org/10.1016/j.scitotenv.2021.145306>, 2021b.

Ma, P. K., Quan, J. N., Dou, Y. J., Pan, Y. B., Liao, Z. H., Cheng, Z. G., Jia, X. C., Wang, Q. Q., Zhan, J. L., Ma, W., Zheng, F. X., Wang, Y. Z., Zhang, Y. S., Hua, C. J., Yan, C., Kulmala, M., Liu, Y. A., Huang, X., Yuan, B., Brown, S. S., and Liu, Y. C.: Regime-Dependence of Nocturnal Nitrate Formation via N<sub>2</sub>O<sub>5</sub> Hydrolysis and Its Implication for Mitigating Nitrate Pollution, *Geophysical Research Letters*. 50, <https://doi.org/10.1029/2023gl106183>, 2023.

Morgan, W. T., Ouyang, B., Allan, J. D., Aruffo, E., Di Carlo, P., Kennedy, O. J., Lowe, D., Flynn, M. J., Rosenberg, P. D., Williams, P. I., Jones, R., McFiggans, G. B., and Coe, H.: Influence of aerosol chemical composition on N<sub>2</sub>O<sub>5</sub> uptake: airborne regional measurements in northwestern Europe, *Atmospheric Chemistry and Physics*. 15, 973-990, <https://doi.org/10.5194/acp-15-973-2015>, 2015.

Niu, Y. B., Zhu, B., He, L. Y., Wang, Z., Lin, X. Y., Tang, M. X., and Huang, X. F.: Fast Nocturnal

314 Heterogeneous Chemistry in a Coastal Background Atmosphere and Its Implications for Daytime  
 315 Photochemistry, *Journal of Geophysical Research-Atmospheres*. 127,  
 316 <https://doi.org/10.1029/2022jd036716>, 2022.  
 317 Shapley, L. S.: A value for n-person games, *Contributions to the Theory of Games*. 1953.  
 318 Sherwen, T., Schmidt, J. A., Evans, M. J., Carpenter, L. J., Grossmann, K., Eastham, S. D., Jacob, D. J.,  
 319 Dix, B., Koenig, T. K., Sinreich, R., Ortega, I., Volkamer, R., Saiz-Lopez, A., Prados-Roman, C., Mahajan,  
 320 A. S., and Ordóñez, C.: Global impacts of tropospheric halogens (Cl, Br, I) on oxidants and composition  
 321 in GEOS-Chem, *Atmospheric Chemistry and Physics*. 16, 12239-12271, [https://doi.org/10.5194/acp-16-](https://doi.org/10.5194/acp-16-12239-2016)  
 322 [12239-2016](https://doi.org/10.5194/acp-16-12239-2016), 2016.  
 323 Thaler, R. D., Mielke, L. H., and Osthoff, H. D.: Quantification of Nitryl Chloride at Part Per Trillion  
 324 Mixing Ratios by Thermal Dissociation Cavity Ring-Down Spectroscopy, *Analytical Chemistry*. 83,  
 325 2761-2766, <https://doi.org/10.1021/ac200055z>, 2011.  
 326 Tham, Y. J., Wang, Z., Li, Q. Y., Yun, H., Wang, W. H., Wang, X. F., Xue, L. K., Lu, K. D., Ma, N., Bohn,  
 327 B., Li, X., Kecorius, S., Gröss, J., Shao, M., Wiedensohler, A., Zhang, Y. H., and Wang, T.: Significant  
 328 concentrations of nitryl chloride sustained in the morning: investigations of the causes and impacts on  
 329 ozone production in a polluted region of northern China, *Atmospheric Chemistry and Physics*. 16, 14959-  
 330 14977, <https://doi.org/10.5194/acp-16-14959-2016>, 2016.  
 331 Wang, H. C., Peng, C., Wang, X., Lou, S. R., Lu, K. D., Gan, G. C., Jia, X. H., Chen, X. R., Chen, J.,  
 332 Wang, H. L., Fan, S. J., Wang, X. M., and Tang, M. J.: N<sub>2</sub>O<sub>5</sub> uptake onto saline mineral dust: a potential  
 333 missing source of tropospheric ClNO<sub>2</sub> in inland China, *Atmospheric Chemistry and Physics*. 22, 1845-  
 334 1859, <https://doi.org/10.5194/acp-22-1845-2022>, 2022a.  
 335 Wang, H. C., Lu, K. D., Guo, S., Wu, Z. J., Shang, D. J., Tan, Z. F., Wang, Y. J., Le Breton, M., Lou, S.  
 336 R., Tang, M. J., Wu, Y. S., Zhu, W. F., Zheng, J., Zeng, L. M., Hallquist, M., Hu, M., and Zhang, Y. H.:  
 337 Efficient N<sub>2</sub>O<sub>5</sub> uptake and NO<sub>3</sub> oxidation in the outflow of urban Beijing, *Atmospheric Chemistry and*  
 338 *Physics*. 18, 9705-9721, <https://doi.org/10.5194/acp-18-9705-2018>, 2018.  
 339 Wang, H. C., Yuan, B., Zheng, E., Zhang, X. X., Wang, J., Lu, K. D., Ye, C. S., Yang, L., Huang, S., Hu,  
 340 W. W., Yang, S. X., Peng, Y. W., Qi, J. P., Wang, S. H., He, X. J., Chen, Y. B., Li, T. G., Wang, W. J.,  
 341 Huangfu, Y. B., Li, X. B., Cai, M. F., Wang, X. M., and Shao, M.: Formation and impacts of nitryl chloride  
 342 in Pearl River Delta, *Atmospheric Chemistry and Physics*. 22, 14837-14858, [https://doi.org/10.5194/acp-](https://doi.org/10.5194/acp-22-14837-2022)  
 343 [22-14837-2022](https://doi.org/10.5194/acp-22-14837-2022), 2022b.  
 344 Wang, Z., Wang, W. H., Tham, Y. J., Li, Q. Y., Wang, H., Wen, L., Wang, X. F., and Wang, T.: Fast  
 345 heterogeneous N<sub>2</sub>O<sub>5</sub> uptake and ClNO<sub>2</sub> production in power plant and industrial plumes observed in the  
 346 nocturnal residual layer over the North China Plain, *Atmospheric Chemistry and Physics*. 17, 12361-  
 347 12378, <https://doi.org/10.5194/acp-17-12361-2017>, 2017.  
 348 Wei, W., Wang, Y., Bai, H. X., Wang, X. Q., Cheng, S. Y., and Wang, L. T.: Insights into atmospheric  
 349 oxidation capacity and its impact on PM<sub>2.5</sub> in megacity Beijing via volatile organic compounds  
 350 measurements, *Atmospheric Research*. 258, <https://doi.org/10.1016/j.atmosres.2021.105632>, 2021.  
 351 Xiong, C., Wang, N., Zhou, L., Yang, F. M., Qiu, Y., Chen, J. H., Han, L., and Li, J. J.: Component  
 352 characteristics and source apportionment of volatile organic compounds during summer and winter in  
 353 downtown Chengdu, southwest China, *Atmospheric Environment*. 258,  
 354 <https://doi.org/10.1016/j.atmosenv.2021.118485>, 2021.  
 355 Yang, C., Dong, H. S., Chen, Y. P., Xu, L. L., Chen, G. J., Fan, X. L., Wang, Y. H., Tham, Y. J., Lin, Z.  
 356 Y., Li, M. R., Hong, Y. W., and Chen, J. S.: New Insights on the Formation of Nucleation Mode Particles  
 357 in a Coastal City Based on a Machine Learning Approach, *Environmental Science & Technology*. 58,

358 1187-1198, <https://doi.org/10.1021/acs.est.3c07042>, 2023.

359 Yun, H., Wang, W. H., Wang, T., Xia, M., Yu, C., Wang, Z., Poon, S. C. N., Yue, D. L., and Zhou, Y.:  
360 Nitrate formation from heterogeneous uptake of dinitrogen pentoxide during a severe winter haze in  
361 southern China, *Atmospheric Chemistry and Physics*. 18, 17515-17527, [https://doi.org/10.5194/acp-18-](https://doi.org/10.5194/acp-18-17515-2018)  
362 [17515-2018](https://doi.org/10.5194/acp-18-17515-2018), 2018.

363 Zare, A., Romer, P. S., Nguyen, T., Keutsch, F. N., Skog, K., and Cohen, R. C.: A comprehensive organic  
364 nitrate chemistry: insights into the lifetime of atmospheric organic nitrates, *Atmospheric Chemistry and*  
365 *Physics*. 18, 15419-15436, <https://doi.org/10.5194/acp-18-15419-2018>, 2018.

366 Zhao, Q. Y., Bi, J., Liu, Q., Ling, Z. H., Shen, G. F., Chen, F., Qiao, Y. Z., Li, C. Y., and Ma, Z. W.:  
367 Sources of volatile organic compounds and policy implications for regional ozone pollution control in an  
368 urban location of Nanjing, East China, *Atmospheric Chemistry and Physics*. 20, 3905-3919,  
369 <https://doi.org/10.5194/acp-20-3905-2020>, 2020.

370