Technical note: Literature based approach to estimate future snow

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Abstract.

The seasonal snow cover in the European Alps is increasingly threatened by rising temperatures due to climate change. Still, downscaled climate projections are lacking for many regions. To address this gap, we developed a literature-based approach for projecting future snow depths, that is applicable to all locations where historical snow depth data is available.

We harmonized heterogeneous literature on future snow depth and snow water equivalent by translating emission scenarios to corresponding temperature scenarios and standardizing seasonal periods. Then, we parameterized localized reduction curves based on elevation, temperature scenarios and local climatologies, as mean snow cover length and mean maximum snow depth. This method was applied to four measurement stations in Switzerland under a $+2°$ C temperature scenario, revealing significant declines in snow depth and season length, especially at lower elevations. Validation against published data shows that the approach captures key trends in snow loss, despite the simplification of climate dynamics.

This resource-efficient method provides a practical tool for estimating climate change related snow depth declines in snow dominated regions, which are lacking highly resolved climate projections, and can support decision-makers in developing adaptation strategies for climate-related challenges.

1 Introduction

Seasonal snow cover plays a crucial role in Alpine hydrology, ecology, and winter tourism. In the context of ongoing climate change, it is increasingly threatened by rising temperatures. Both, snow depth and snow water equivalent (SWE) have shown substantial decreases across the European Alps in recent decades (Marty et al., 2025; Ranzi et al., 2024; Broust et al., 2024), with future projections indicating further declines across all ranges of elevations and regions (Kotlarski et al., 2022; Bülow et al., 2025). While detailed climate projections exist for selected study areas and provide detailed insights into future snow cover (Marty et al., 2017; Schmucki et al., 2015; Willibald et al., 2020; Verfaillie et al., 2018), many Alpine regions still lack high resolution projections. This presents a challenge for practitioners and decision-makers requiring localized snow cover information to support climate adaptation strategies.

To bridge this gap within a short-term project with limited resources, we developed a transferable, literature-based approach, which is applicable to climatological datasets (both point-based and gridded) by synthesizing existing studies to estimate future snow cover changes.

2 Methods

2.1 Synthesizing literature values

The heterogeneity of available studies posed several methodological challenges, e.g. the heterogeneity in projected regions and elevations, emission scenarios and examined variables.

30   Reported variables ranged from decreases in seasonal means for different period (Willibald et al., 2020; Kotlarski et al., 2022; Morin et al., 2018; Marty et al., 2017; Verfaillie et al., 2018), season lengths for different thresholds in snow depths (Willibald et al., 2021; Verfaillie et al., 2018; Morin et al., 2018), monthly values (Bülow et al., 2025; Marty et al., 2017) or seasonal evolutions (Schmucki et al., 2017; Schmucki, 2015; Fiddes et al., 2022) of either SWE or snow depths. Since the relative reductions in SWE and snow depth were comparable across these studies (Schmucki et al., 2015; Verfaillie et al.,

35 2018), we treated the reduction values as interchangeable.

    Tables A1 and A2 in the appendix summarize the examined variables. These studies were generally divided into two groups: Literature-Fit and Literature-Validation (see comments in Tables A1 and A2). The data in Literature-Fit provided daily or monthly snow depths, both for the reference period and the future projection, which could be used to train the reduction curves in Section 2.2. The data in Literature-Validation reported seasonal reduction values or decreases in season length and was used

40 for validation only (Sections 2.4 and 3.3).

2.1.1 Translating emission scenarios to temperature scenarios

Most studies referred to different emission scenarios (e.g., various RCPs). To address this complexity, we standardized reported climate change scenarios by translating all RCPs, reference periods, and projected periods into corresponding temperature scenarios, if temperature scenario was not directly reported. To this end, we used the reports CH2011 and CH2018 (CH2018,

45 2018; CH2011, 2011) and summarized emission and corresponding temperature scenarios in Table B1 (see Appendix).

2.1.2 Synthesizing different seasonal means

Reduction values for seasonal means in the Literature-Validation dataset were reported for various time periods, ranging from the shortest—December to February (DJF, 3 months)—to the longest—September to August (SONDJFMAMJJA, 12 months). As reduction values may highly vary depending on the period, as illustrated in an example below, literature values are difficult

50 to compare. For consistency, we synthesized all results to a common six-month period from November to April (NDJFMA).

    To this end, we used the Literature-Fit dataset, which was linearly interpolated to daily values, to calculate mean snow depths for both the reference period (HS_{ref}) and the future projection (HS_{fut}) across different seasonal windows (e.g., NDJFMA). The relative decrease in mean snow depth for NDJFMA was then calculated as:

$$\text{NDJFMA-decrease} = \frac{HS_{fut}(\text{NDJFMA}) - HS_{ref}(\text{NDJFMA})}{HS_{ref}(\text{NDJFMA})} \qquad (1)$$
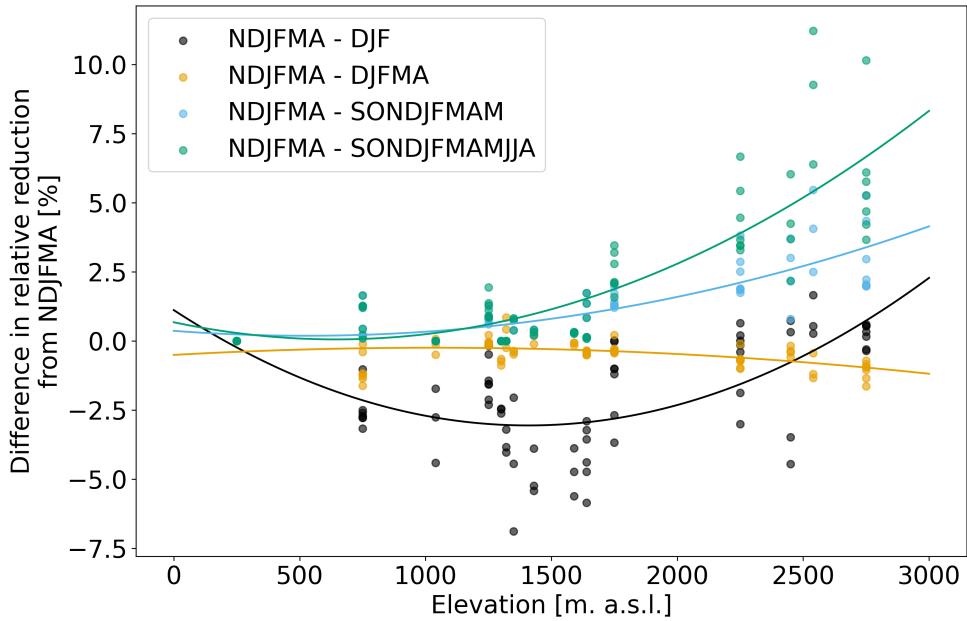
Figure 1. Relationship of relative decreases in mean snow depths for different periods (colors) compared to NDJFMA period. Values were derived from studies which provided monthly or daily snow depths and decreases. For each period we computed a second order polynomial fit with elevation.

Reduction values for other periods were computed analogously and compared to the NDJFMA-decrease (Figure 1). For illustration, we use the example of the lower most point at 1350 m a.s.l. in Figure 1, which comes from Schmucki (2015) (see Figure 2): The mean reference snow depth during the NDJFMA-period is $HS_{ref}(\text{NDJFMA}) = 68$cm, whereas the future mean snow depth is $HS_{fut}(\text{NDJFMA}) = 25$cm, corresponding to a NDJFMA-decrease of -63%. In comparison, the DJF-decrease is only -56% ($HS_{ref}(\text{DJF}) = 78$cm, $HS_{fut}(\text{DJF}) = 34$cm). Thus, if the DJF-decrease were used in place of the NDJFMA-decrease, the relative reduction would be underestimated by about -7%. These differences in relative reduction between the NDJFMA and DJF periods (in this example -7%) and other periods are shown for the full Literature-Fit dataset in Figure 1.

Depending on the specific period and elevation, deviations of up to 10% from the NDJFMA-decrease were observed. To account for these variations, we applied a second-order polynomial fit to adjust reduction values based on both period length and elevation (lines in Figure 1).

2.2 Fitting of Seasonal Reduction Curves

We analyzed the seasonal snow depth reduction curves for the Literature-Fit data, by first linearly interpolating reported values to daily data. If reduction values were not provided, we computed those from reference and future snow depth or SWE data.

Figure 2 (top) shows reference and future snow depth at 1350 m a.s.l. under a +2.4 °C scenario from Schmucki (2015). Peak snow depth decreased from 116 cm to 52 cm, with a reported reduction in season length of approximately 20 days. In most of the Literature-Fit data this shortening is asymmetric: for example, Figure 2 (top) indicates a slight delay of a few days in future accumulation, whereas ablation occurs approximately two weeks earlier. In contrast, the corresponding future relative snow depths (Fig. 2, bottom), which are equivalent to relative reduction curves, appear symmetric around the value b. The reduction curve peaked shortly before the seasonal maximum ($b = 144$ DOWY), and approached 0 % toward both ends of the season. This shape implies both a delayed onset and an earlier end of the season, reflecting this general shortening of the snow duration. All relative reduction curves consistently showed this behavior without any systematic asymmetry, therefore, reduction curves $f_{red}(x)$ could be well approximated by a quadratic function:

$$f_{red}(x) = -100 + a - \frac{a}{c^2} \cdot (x - b)^2, \quad \text{with } f(x) \geq -100, \tag{2}$$

To express the relative future snow depth $f(x) = f_{red}(x) + 100$ directly, this can be rewritten as:

$$f(x) = a - \frac{a}{c^2} \cdot (x - b)^2, \quad \text{with } f(x) \geq 0, \tag{3}$$

where:

- x is the day of water year (DOWY, ranging from 1 to 366, with DOWY 1 = 1. September),

- a denotes the maximum of $f(x)$, corresponding to the highest relative snow depth in future,

- b is the day of the year (DOWY) on which this maximum occurs,

- c defines half the width of the curve, and thus approximates half the snow season length (i.e., days with snow on the ground before and after b).

Finding trends for b and c directly, would imply to fix the position of the peak of the reduction curve to a specific day in the season (b) and fix the length of snow duration to a certain number of days (c). However, snow duration can highly differ between regions and locations. To preserve the local climatology, instead of finding trends for b and c, we explored trends for the relative snow duration length (Δc) and the position of the peak of the reduction curve relative to the peak snow depth (Δb):

$$\Delta b = b - \text{DOWY}(HS_{ref,max}) \text{ [days]} \tag{4}$$

$$\Delta c = \frac{len(HS_{fut} > 0\text{cm})}{len(HS_{ref} > 0\text{cm})} \cdot 100 \text{ [%]} \tag{5}$$
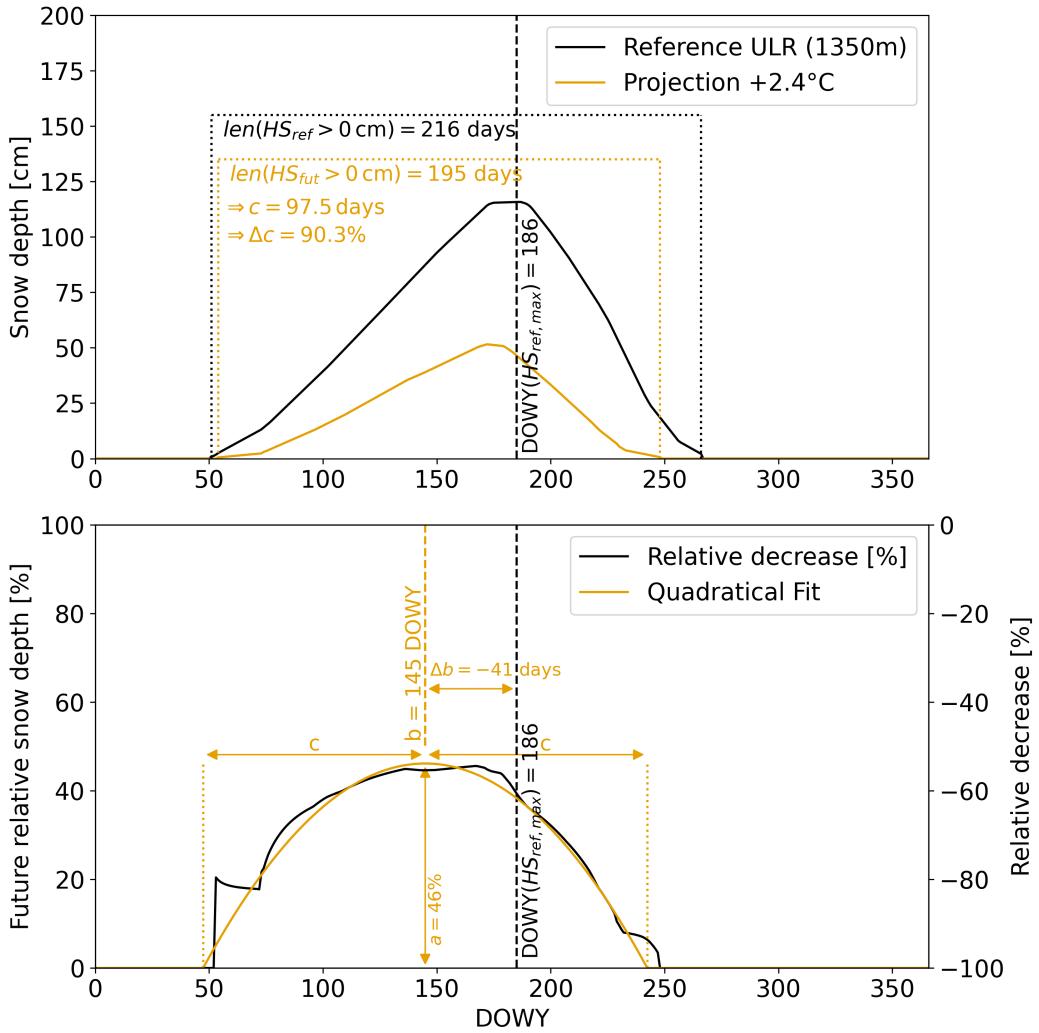
Figure 2. (Upper) Reference and future snow depth from Schmucki (2015) (temperature scenario: $+2.4\,°\mathrm{C}$, elevation: $1350\,\mathrm{m}$ a.s.l., monthly averages) and (lower) corresponding future relative snow depths. The fitting parameter a and Δb are marked in orange, where a corresponds to the maximum future relative snow depth, Δb is the position of the peak of the reduction curve relative to the peak in reference snow depth, and Δc is the fraction of future season length to reference season length, which are also marked (upper). Orange line shows the quadratic fit for the reduction curve. x-axis is the day of water year (DOWY) starting on 1. September.

In Figure 2, the maximum future snow depth $a = 47\%$ and the reduction curve peaks on $b = 144$ DOWY (corresponding to 23. January), hence the peak of the reduction curve is 41 days prior to the peak in reference snow depth ($\Delta b = -41$). The total snow duration decreases from 216 days to 195 days, resulting in a relative decrease of $\Delta c = 90\%$.

Finally, we compared all reduction parameter (a, Δb, and Δc) from the Literature-Fit data with temperature scenario and elevation (see Section 3.1) and trained a linear regression model. To avoid overfitting, we used linear terms of temperature scenario and elevation and their interaction terms (elevation $\times \Delta T$), which were scaled before training. While the reduction parameters a, Δb and Δc can be computed for any given elevation and temperature scenario, these parameters were trained on data with elevations ranging from 750 m and 2750 m and temperature scenarios from $+1.1°$C to $+4.8°$C and should be treated with caution outside these ranges.

2.3 Apply reduction curve to project future snow evolution

We applied the reduction curves to retrieve future snow depths for four measurement stations in Switzerland: Weissfluhjoch (WJ, 2540 m a.s.l.), in the eastern Swiss Alps, Maloja (MA, 1810 m a.s.l.), in the southern Swiss Alps, Saanenmöser (SM, 1390 m a.s.l.), in the western Swiss Alps, and Engelberg (EN, 1023 m a.s.l.), in the central Swiss Alps. These stations provide daily manually measured snow depth data from winter season 1991-2020 (30 years). We computed daily median, as well as the 5th and the 95th percentile for those stations, as our reduction curves were trained on data, which does not account for extreme events.

We chose a temperature scenario of $\Delta T = +2°$C for the projections then computed the reduction parameters a, Δb and Δc for the given elevations. As our reference period 1991-2020 experienced a mean annual temperature increase of $+0.5°$C compared to the period 1981-2010 (Senoner et al., 2023), this temperature scenario of $\Delta T = +2°$C refers to the climate period 2043-2072 for the RCP8.5 scenario of CH2018 (2018).

Climate projections show uncertainties in temperature scenarios of around $+/-1°$C for each RCP-scenario between years 2000 and 2100 (CH2018, 2018). Therefore, we applied the same uncertainty range for the snow projections as follows:

- Temperature scenario ΔT for median snow depths (here: $\Delta T = +2°$C)

- "$\Delta T + 1°$C " scenario used for 5th percentile (higher increase, more pessimistic scenario, here: $+3°$C)

- "$\Delta T - 1°$C" scenario used for 95th percentile (less increase, more optimistic scenario, here: $+1°$C)

To apply the reduction curve (Equation 3) to climatological evolutions we have to compute b and c from Δb and Δc, respectively. To this end, we first determined $\text{DOWY}(HS_{ref,max})$ and the reference season length $len(HS_{ref} > 0\text{cm})$. We suggest to smooth reference snow evolutions prior to determining $\text{DOWY}(HS_{ref,max})$ using a running mean of 30 days, to smoothen temporal variability in the data. Then b and c can be computed as followed:

$$b = \text{DOWY}(HS_{ref,max}) + \Delta b \tag{6}$$

$$c = \frac{len(HS_{ref} > 0\text{cm}) \times \Delta c}{2} = \frac{len(HS_{fut} > 0\text{cm})}{2} \tag{7}$$

Finally, using the reduction curve (Equation 3) and reference snow depths $HS(x)_{ref}$, future snow depths $HS(x)_{fut}$ can be computed as follows:

$$HS(x)_{fut} = f(x) \cdot HS(x)_{ref} \qquad (8)$$

2.4 Variables used for validation

130 After synthesizing all seasonal decreases from the Literature-Fit and Literature-Validation dataset into NDJFMA-decreases (Section 2.1.2), we compared these Literature values to the NDJFMA-decreases from our projections of the four stations using Equation 1.

Furthermore, we also looked at the relative reduction in season lengths by counting the days for which a certain snow depth was reached, both in the reference $len(HS_{ref} > x)$ and for the projections $len(HS_{fut} > x)$. The relative reduction in season

135 length $(HS > x)$-decrease was then computed as follows:

$$(HS > x)\text{-decrease} = \frac{len(HS_{fut} > x) - len(HS_{ref} > x)}{len(HS_{ref} > x)} \cdot 100 \qquad (9)$$

The Literature-Validation dataset contains decreases in season lengths for the following thresholds: $> 5\,\mathrm{cm}$, $> 30\,\mathrm{cm}$, $> 50\,\mathrm{cm}$ and $> 100\,\mathrm{cm}$. As most values were reported for the former two thresholds, we chose the following variables for validation:

140 – Relative decrease in mean November-April snow depth (NDJFMA-decrease).

– Relative decrease in season length with more than $30\,\mathrm{cm}$ snow on the ground $(HS > 30\mathrm{cm})$-decrease.

– Relative decrease in season length with more than $5\,\mathrm{cm}$ snow on the ground $(HS > 5\mathrm{cm})$-decrease.

Projected relative decreases were calculated for the median snow depths and 5-95 percentiles.

3 Results

145 ### 3.1 Reduction curve parameters with respect to temperature and elevation

Figure 3 (left) presents the reduction parameter a, which corresponds to the maximum future relative snow depth. As expected, a decreases with increasing temperature scenarios, indicating less snow under warmer conditions. Furthermore, a increases with elevation, indicating that the decrease of future snow depth is more pronounced at lower elevations.

Although no consistent trend was observed for the parameter b itself, Δb were predominantly negative (Figure 3, middle),

150 suggesting a shift in the timing of peak snow depth towards earlier in the season. Notably, this shift becomes more pronounced with elevation, as Δb decreases with higher elevation. Δb positions the peak of the reduction curves relative to peak in snow depth, which peaks later in the season for higher elevations as the accumulation period is longer.
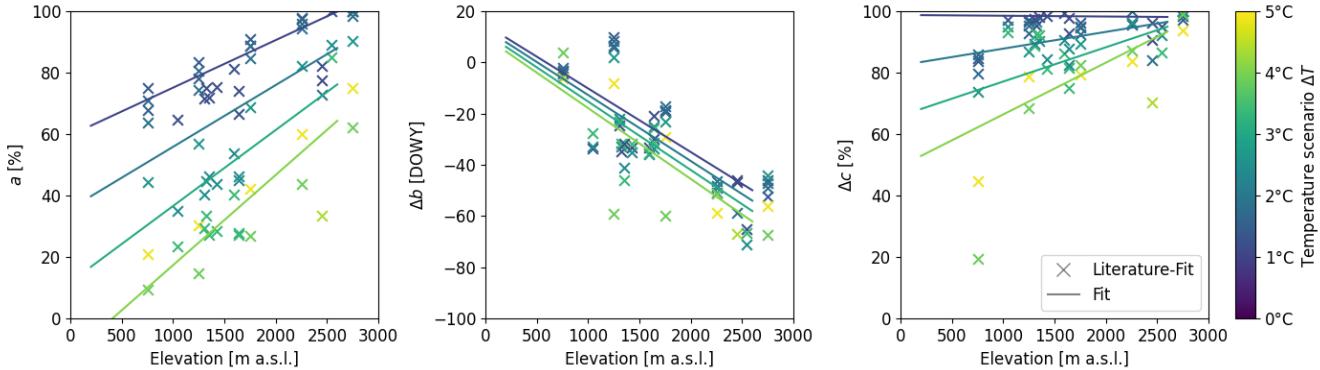
Figure 3. Reduction parameters from the Literature-Fit dataset and corresponding fits with elevation and temperature scenario, which were used for computing reduction curves for future snow projections: (left) maximum of the future relative snow depth a, (middle) days between peak reduction and peak snow depth Δb, and (right) the future relative season length Δc.

The relative change in season length Δc is shown in Figure 3 (right). All studies indicated shorter snow seasons in future scenarios, with the reduction in season length being more substantial at lower elevations and under higher temperature scenarios.

The following regression formulas were derived to describe the relationships of the reduction parameters with temperature and elevation (lines in Figure 3):

$$a = 83.51239 - 23.89164 \cdot \Delta T + 0.01085 \cdot h + 0.00463 \cdot \Delta T \cdot h, \quad \text{with } a \in [0, 100] \tag{10}$$

$$\Delta b = 16.28888 - 1.61312 \cdot \Delta T - 0.02390 \cdot h - 0.00094 \cdot \Delta T \cdot h \tag{11}$$

$$\Delta c = 115.18564 - 16.41657 \cdot \Delta T - 0.00595 \cdot h + 0.00570 \cdot \Delta T \cdot h \tag{12}$$

3.2 Projected snow evolutions

Snow projections under a $\Delta T = +2°C$ temperature scenario indicate a decline in snow depth across all sites (Figure 4). At Weissfluhjoch, the peak median snow depth decreases from 215 cm to 171 cm, while in Saanenmöser it drops from 64 cm to 36 cm. In Engelberg, the median snow depth during the reference period never exceeded 30 cm, and in 5% of winters, snow depth remained at 0 cm throughout the entire year.

All projections also indicate shorter snow seasons in the future. For example, at Weissfluhjoch, the snow season (with $HS > 0$ cm) is projected to begin approximately two weeks later and to end nearly two weeks earlier on average.
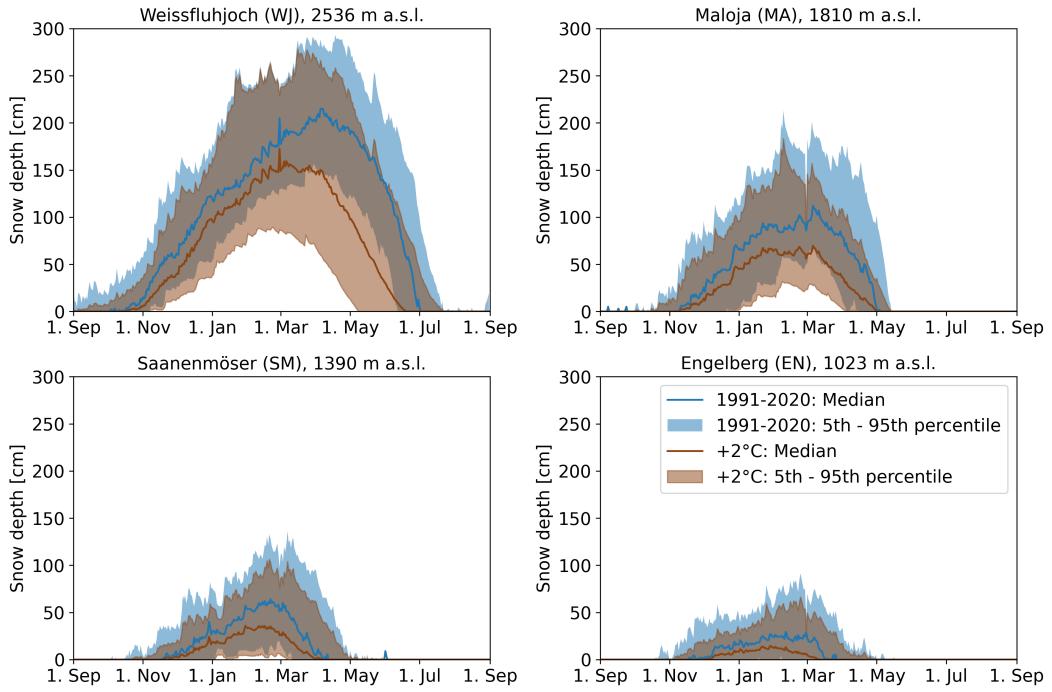
Figure 4. Snow depth evolution for Weissfluhjoch (upper left), Saanenmöser (upper right), Maloja (lower left) and Engelberg (lower right) for the reference period 1991-2020 and a future projection for the temperature scenario $+2°C$.

3.3 Validation of snow projections

We validated the projections by comparing the projected decreases at all four study sites with values reported in the literature (Figure 5). Projected decreases shown as bars represent the 5th and 95th percentiles, while black lines indicate the decrease in median snow depth. Since literature values were not always based on the exact same temperature scenarios or elevations, comparisons were quantitatively made based on the range of values rather than exact matches.

Both the Literature-Fit and Literature-Validation datasets show similar trends (percentage changes) for the NDJFMA-decrease, the (HS> 30 cm)-decrease and the (HS> 5 cm)-decrease: the relative decreases are stronger with higher temperature scenarios and at lower elevations (Figure 5). The projections align well with the reported ranges and replicate the expected elevation-dependent trends. Both, literature data and projections show weaker (HS> 5 cm)-decreases compared to the (HS> 30 cm)-decrease.

Uncertainty ranges in the projections, showing the relative decrease in the 5th and the 95th percentile, seemed to be occasionally larger than uncertainty ranges from literature values. This is due to the methodology (see Section 2.3), as we are projecting median and 5-95 percentiles rather then individual years: For instance, the 5th percentile snow depth at Engelberg (EN) is zero at each calendar day in the reference period (see Figure 4). This indicates that there is currently not one period
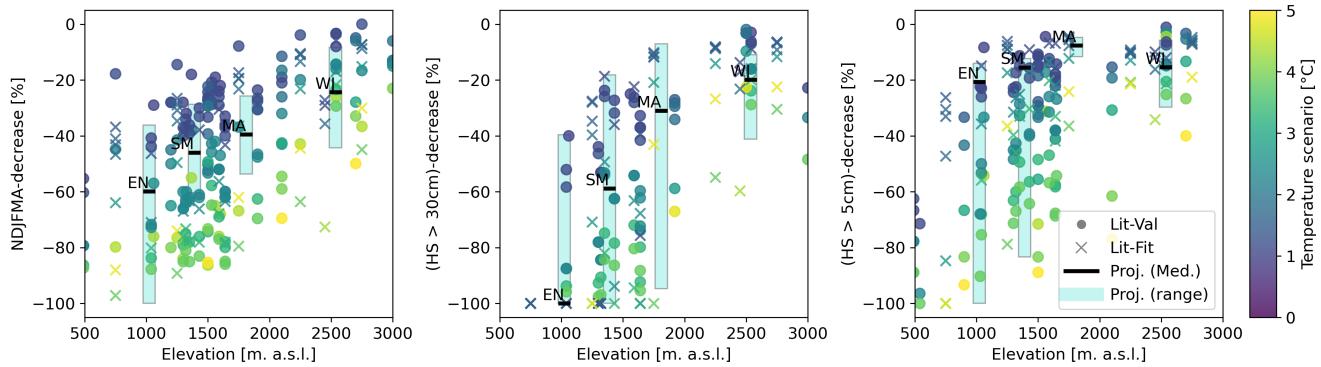
Figure 5. Projected (bars) and literature based (markers) values for (left) the relative decrease in mean November to April snow depths NDJFMA-decrease, (middle) relative decrease in season length with more than 30 cm snow (HS$>$ 30 cm)-decrease and (right) decrease in season length with more than 5 cm snow (HS$>$ 5 cm)-decrease with elevation. Colors indicate temperature scenarios. Projections show median decrease (black line) and 5th and 95th percentile (bars) for the four stations Weissfluhjoch (WJ), Maloja (MA), Saanenmöser (SM), and Engelberg (EN).

(day) throughout a year, where snow on the ground can be guaranteed in Engelberg. As such the projected decrease in Figure 5
185   was set to -100%. On the other hand, we want to highlight that this does not imply that 5% of the future winters will be entirely snow-free.

4   Discussion

This study presents a practical and efficient alternative to high-resolution climate model downscaling for estimating future changes in snow depth and season length. Synthesizing existing literature and applying a simplified yet structured method to
190   observational snow data allows to assess climate impacts in regions where localized projections are scarce or unavailable.

A major strength of this method lies in its low resource demand. No computational complex physically-based snow models or dynamically downscaled regional climate simulations are required. Instead, observed snow data in combination with parameterized reduction curves is sufficient to obtain snow projections for different climate scenarios. As such, the approach inherently captures local climatology and site-specific features (e.g., exposure, shading, cold pool valleys) that are often missed
195   in coarse-grid climate models (Frei, 2014). This is especially valuable for metrics like snow season length above a specific threshold (e.g., $HS > 30$ cm), which are highly sensitive to local topography and micro-climatic conditions.

To synthesize existing literature on future snow depth reductions, we translated climate scenarios, reference periods, and projection periods into temperature scenarios. This approach offers several advantages. First, it enables to directly compare literature results and helps to quantify uncertainties across different studies. Second, it substantially improves communication
200   with policy makers, as many national and international climate targets are framed around temperature thresholds. This com-

munication will make the results more accessible to non-scientific audiences. Finally, this methodology can be extended and updated as new studies based on the latest climate scenarios become available.

Despite the above described strengths, our simple approach has several important limitations, in the first place the simplification of climate dynamics. While temperature change is the dominant driver of future snow loss in the Alps (Marty et al., 2017; Verfaillie et al., 2018; Kotlarski et al., 2022), other climatological uncertainties such as changing precipitation, or change of large scale weather patterns are not explicitly incorporated. These factors are only indirectly captured, as the reduction parameters were trained on projections that do consider such dynamics. Still, this indirect inclusion does not allow assessing their future contributions. Other studies performed sensitivity analyses (e.g., adding $2\,^{\circ}\mathrm{C}$ and $\pm 10\%$ changes in precipitation as in López-Moreno et al. (2020); Richter et al. (2020)) to compare the impact on the future snow cover.

Furthermore, the quadratic approximation of snow reduction curves and the linear interpolation of reduction parameters with temperature and elevation may not capture the full complexity of snow accumulation and melt processes. These simplifications may limit accuracy in regions with strong seasonal asymmetries or where snow dynamics are highly variable.

The reduction parameters where trained for elevations ranging between $750\,\mathrm{m} - 2750\,\mathrm{m}$ and temperature scenarios between $+1.1^{\circ}\,\mathrm{C}$ and $+4.8^{\circ}\mathrm{C}$. Extrapolations outside these ranges should be treated with caution.

The literature-derived values often lack clear definitions of ensemble spreads, e.g. spread of mean of individual ensemble members (Schmucki et al., 2015) or spread of all modeled snow depth evolutions. As a robust statistical assessment was not possible, we adopted a simple $+/-1^{\circ}\mathrm{C}$ variation around the central temperature scenario to approximate uncertainty bounds (5th and 95th percentiles). Although this is a coarse approximation, the resulting projection spread remains within the variability reported across different studies and models (Marty et al., 2017; Bülow et al., 2025).

However, while the approach offers high transferability and ease of use, it simplifies complex climatic and geographic interactions. Thus, it is best used as a first-order estimate in regions where detailed projections are not available, or to complement more detailed modeling efforts.

5   Conclusion

This study introduces a resource-efficient approach to project future snow cover evolution across Alpine regions. Literature values were used to obtain localized reduction curves depending on temperature scenario and elevation. The relative reductions curves could be described with three metrics: (1) maximum relative future snow depth, (2) shift in the timing of peak snow depth, and (3) relative shortening of the snow season. Projections show consistent and plausible snow depth trends for different elevations and temperature scenarios.

This method is rather simplified and rudimentary, still it offers a robust and adaptable framework for estimating future local or regional snow depth changes in the absence of high-resolution climate projections. As such, it provides a valuable tool for the assessment of climate impacts and development of adaptive strategies in snow-dependent regions of the European Alps.

Code and data availability. Literature data and projected snow depths can be downloaded here: https://www.doi.org/10.16904/envidat.719. In-situ snow depth data from SLF stations can be freely downloaded from: https://www.slf.ch/en/services-and-products/slf-data-service.

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235 **Appendix A: Table of Literature values used for training and validation**

Literature	Reports Reference period	Scenario	Climate period	Elevation a.s.l. [m]	Region	Comments
Bülow et al. (2025)*	CH2018 1971-2000	RCP26 RCP45 RCP85	2021-2050 2069-2098	0-500 500-1000 1000-1500 1500-2000 2000-2500 2500-3000	Alps	Literature-Fit monthly values
Marty et al. (2017)	CH2011 1980–2009	A2	2071-2100	0-500 500-1000 1000-1500 1500-2000 2000-2500 2500-3000	Region Aare	Literature-Fit monthly values
Schmucki et al. (2017) Schmucki (2015)	CH2011 1984–2010	A1B	2020-2049 2045-2074 2070-2099	2540 1640 1640 1590 1430 1350 1320 1300 1040	Weissfluhjoch San Bernardino Zermatt Davos Montana Ulrichen Adelboden Scuol Engelberg	Literature-Fit daily values
Fiddes et al. (2022)	CH2018 1981–2010	RCP2.6 RCP8.5	2031-2060 2070-2099	2450 m	IMIS	Literature-Fit daily values
Willibald et al. (2020)	CH2018 1980-2009	RCP8.5	2010-2039 2040-2069 2070-2099	2540 1640 1590 1430 1350 1320 1300 1040	Weissfluhjoch Zermatt Davos Montana Ulrichen Adelboden Scuol Engelberg	Literature-Validation Nov-Apr
Willibald et al. (2021)	CH2018 1980-2009	RCP8.5	2010-2039 2040-2069 2070-2099	2540 1590 1040	Weissfluhjoch Davos Engelberg	Literature-Validation season length (HS> 30cm)
Verfaillie et al. (2018)	CH2018 1986–2005	RCP2.6 RCP4.5 RCP8.5	2022-2038 2042-2058 2062-2078 2082-2098	1500	Chatreuse	Literature-Validation Dec-Apr season length (HS> 50 cm)

Table A1. Overview over Literature values, which were used for training (Literature-Fit, first four studies in Table A1) and validation (Literature-Validation). Studies which reported values for SWE were marked with *. (Part 1, part 2 see Table A2)

Literature	Reports Reference period	Scenario	Climate period	Elevation a.s.l. [m]	Region	Comments
Schmucki et al. (2015)	CH2011 1984–2010	A1B A2	2020-2049 2045-2074 2070-2099	2540 1640 1640 1590 1430 1350 1320 1300 1040	Weissfluhjoch San Bernardino Zermatt Davos Montana Ulrichen Adelboden Scuol Engelberg	Literature-Validation Dec-Feb season length (HS> 30 cm)
Kotlarski et al. (2022)*	CH2018 1981-2010	RCP2.6 RCP4.5 RCP8.5	2070-2099	0-500 500-1000 1000-1500 1500-2000 2000-2500 2500-3000	Alps	Literature-Validation Sep-May
Kotlarski et al. (2022)	CH2018 1981-2010	RCP2.6 RCP4.5 RCP8.5	2021–2094	1200 2100 3000	Mont-Blanc	Literature-Validation Nov-Apr
Kotlarski et al. (2022)	CH2018 1971-2000	RCP2.6 RCP4.5 RCP8.5	2021–2050 2070–2099	1920 2500 3000	Ötztaler Alps	Literature-Validation season length (HS> 30cm)
Marty et al. (2017)	CH2011 1999-2012	A1B A2 RCP3PD	2020–2049 2045–2074 2070–2099	1530 1903	Aare Graubünden	Literature-Validation Sep-Aug
Marty et al. (2017)	CH2011 1999-2012	A2	2020–2049 2045–2074 2070–2099	540 1030 1650	Bern Grindelwald Mürren	Literature-Validation season length (HS> 5cm)
Marty et al. (2017)	CH2011 1999-2012	A2	2020–2049 2045–2074 2070–2099	3000	Aare	Literature-Validation season length (HS> 30cm)
Morin et al. (2018)	IPCC2013 1986-2005	RCP2.6 RCP4.5 RCP8.5	+1.5°C +2°C +3°C +4°C +5°C	1500 2100 2700	Mont-Blanc	Literature-Validation Dez-Apr
Morin et al. (2018)	IPCC2013 1986-2005	RCP2.6 RCP4.5 RCP8.5	+1.5°C +2°C +3°C +4°C +5°C	900 1500 2100 2700	Pyrenees	Literature-Validation season length (HS> 5cm)

Table A2. Overview over Literature values, which were used for training (Literature-Fit, first four studies in Table A1) and validation (Literature-Validation). Studies which reported values for SWE were marked with *. (Part 2, part 1 see Table A1)

Appendix B: Table: Linking RCP-Scenarios and Climate periods to temperature scenarios

Author contributions. BR: Data analysis, Conceptualization, Methodology, Software, Writing – original draft. CM: Conceptualization, Methodology, Writing – review & editing.

Competing interests. The authors declare that they have no conflict of interest.

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Table B1. Mapping of emission scenario, reference and climate period to temperature scenario (ΔT).

Report	Emission scenario	Reference period	climate period	ΔT [°C]
CH2011	A1B	1984–2010 1999–2012	2020–2049	+1.2
			2045–2074	+2.4
			2070–2099	+3.3
	A2	1984–2010 1999–2012 1980–2009	2020–2049	+1.1
			2045–2074	+2.3
			2070–2099	+3.8
			2071–2100	+3.8
	RCP3PD	1999–2012	2020–2049	+1.2
			2045–2074	+1.2
			2070–2099	+2.4
CH2018	RCP2.6	1971–2000	2021–2050	+1.5
			2069–2098	+1.6
		1981–2010	2031–2060	+1.1
			2070–2099	+1.2
		1986–2005	2024–2037	+0.5
			2044–2057	+0.8
			2064–2077	+0.9
			2084–2097	+0.8
	RCP4.5	1971–2000	2021–2050	+1.7
			2069–2098	+2.7
		1986–2005	2024–2037	+0.8
			2044–2057	+1.4
			2064–2077	+1.8
			2084–2097	+1.9
	RCP8.5	1971–2000	2021–2050	+1.8
			2069–2098	+4.8
		1981–2010	2031–2060	+1.9
			2070–2099	+4.4
		1980–2009	2010–2039	+1.0
			2040–2069	+2.4
			2070–2099	+4.3
		1986–2005	2024–2037	+0.8
			2044–2057	+1.7
			2064–2077	+3.0
			2084–2097	+4.1

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