

Intercomparison of tropopause height climatologies: High-Resolution radiosonde measurements versus ERA5 reanalysis

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Abstract. The tropopause plays a critical role in stratosphere-troposphere exchange and climate change. Its height is widely defined based on the World Meteorological Organization (WMO) threshold temperature gradient. High-resolution (5–10 m) soundings, therefore, are expected to substantially minimize uncertainties of tropopause height (TH) arising from limited vertical resolution and imprecise temperature measurements. The high-resolution radiosonde data, accumulated from 2000 to 2023 from a globally distributed, sparse network (about 1.5 million profiles from 222 stations), offers valuable insights into climatological tropopause variability. While radiosonde observations are limited by spatiotemporal coverage, European Centre for Medium-Range Weather Forecasts Reanalysis v5 (ERA5) reanalysis datasets offer globally complete tropopause representations. To leverage both the high resolution of radiosonde measurements and the global coverage of ERA5, this study compares their TH estimates and analyzes long-term trends across different latitude zones and seasons. The results indicate that the mean and absolute differences (radiosonde minus ERA5) in TH were 32 m and 336 m, respectively, with larger discrepancies observed during the spring season in the tropics ($\pm 20^\circ$). Overall, point-to-point comparisons indicate that ERA5 effectively captures climatological TH variations in both time and space. Long-term trend analyses revealed increases of +9 m/year (radiosonde) and +7 m/year (ERA5) based on point-to-point comparisons. However, these site-specific trends may differ substantially from the long-term trends observed in ERA5 with complete spatiotemporal resolution, even showing opposite trends. Therefore, continued accumulation of high-resolution radiosonde profile data is crucial to further characterize tropopause changes in a warming climate.

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Introduction

The tropopause, marking the boundary between the turbulent troposphere and the stably stratified stratosphere, is a “gate” for exchange of energy, air masses, water vapor and short-lived substances (Fueglistaler, 2009). Furthermore, given the impact of global warming and ozone depletion on the troposphere and stratosphere, tropopause variations can serve as an indicator of anthropogenic environmental influences (Santer et al., 2003). Moreover, its extreme sensitivity to climate variability and change makes the tropopause a pivotal factor in understanding and predicting future climate scenarios (Sausen and Santer, 2003; Seidel and Randel, 2006).

The tropopause exhibits unique chemical and dynamical characteristics, and its maintenance relies on complex interactions between large-scale and small-scale circulation patterns, deep convection, cloud formation, and radiation (Randel and Jensen, 2013). For instance, water vapor abundance can influence tropopause height (TH), as an increase in TH accompanies increased optical thickness to maintain a constant emitted temperature. In addition, TH is generally controlled by a combination of diabatic forcing and adiabatic dynamical effects (Zurita-Gotor and Vallis, 2013). These intricate mechanisms eventually lead to a marked difference in TH between the tropics and the poles, with the annual average reaching approximately 16 km in the tropics and 8 km in the polar regions. The subtropical jet streams (STJs), which are typically located at 20°–40° latitude in each hemisphere and 12 km height (Manney and Hegglin, 2018) will reinforce the Hadley circulation via “eddy pump” (Staten et al., 2018). The large-scale downwelling in the subtropical Hadley circulation sharply lowers the tropopause, sometimes creating a discontinuity known as the “subtropical tropopause break”, which aligns with the STJ (Turhal et al, 2024).

There are multiple ways to calculate the TH, relying on several empirical criteria based on properties exhibiting sharp transitions between the troposphere and stratosphere. The cold point tropopause (CPT) gets the tropopause via the minimum temperature in the vertical temperature profile, which is unsuitable for extratropical regions (Highwood and Hoskins, 1998), and the dynamic tropopause (WMO, 1985; Hoinka, 1998), typically defined by potential vorticity (PV) thresholds of 1.5–4 potential vorticity unit (PVU) (Turhal et al., 2024), is less reliable in regions of low absolute potential vorticity, such as the tropics, and sometimes in mid-latitudes where strong anticyclonic flow prevails (Hoerling et al., 1991). However, the tropopause definition (WMO, 1957), which is thermodynamic, proposed by the World Meteorological Organization offers a more robust global approach (though it may occasionally fail in polar regions), providing reliable TH estimates from various datasets (Hoffmann and Spang, 2022).

Current primary data sources for TH determination include in-situ measurements from radiosonde, meteorological reanalysis datasets, and radio occultation data from the Constellation Observing System for Meteorology Ionosphere and Climate (COSMIC) Data Analysis and Archive Center (CDAAC) of the Global Navigation Satellite System (GNSS). Radiosonde data, collected at stations worldwide, feature long-term records with high accuracy and reliability, yet suffer from sparse and highly uneven global coverage. Global atmospheric reanalysis products provide comprehensive long-term atmospheric information through continuously improved forecast models, observational data, and assimilation

schemes (Fujiwara et al., 2017). Modern reanalysis datasets, such as the state-of-the-art ERA5 developed by the European Centre for Medium-Range Weather Forecasts (ECMWF), and Modern-Era Retrospective analysis for Research and Applications, Version 2 (MERRA-2) produced by National Aeronautics and Space Administration (NASA) Global Modeling and Assimilation Office (GMAO) (Gelaro et al., 2017), overcome the spatial and temporal limitations of observational records, offering a global perspective of the tropopause. GNSS-based datasets (e.g., COSMIC and the Radio Occultation Meteorology Satellite Application Facility (ROM SAF)) provide high-density measurements with near-global coverage, making them particularly suitable for analyzing global-scale tropopause characteristics (Son et al., 2011).

More and more evidence suggests an upward trend in TH under a changing climate (Santer et al., 2003; Sausen and Santer, 2003; Seidel and Randel, 2006; Añel et al., 2006). For instance, Seidel and Randel (2006) observed a 64 ± 21 m/decade upward trend from 1980 to 2004, and Son et al. (2009) projected continued future increase, albeit with a weaker trend. Analyses by Xian and Homeyer (2019) using radiosonde observations and reanalysis datasets from 1981–2015 detected a significant upward trend (40–120 m/decade). Meny et al. (2021) reported a TH increase of 50–60 m/decade (2001–2020) in the Northern Hemisphere based on radiosonde data. More recently, Zou et al. (2023) leveraged European Centre for Medium-Range Weather Forecasts (ECMWF) Reanalysis v5 (ERA5) data to reveal a widespread upward and cooling trend in the tropical tropopause from 1980–2021, demonstrating an increase of approximately 60 ± 10 m/decade (95% confidence). It is worthwhile to note that these radiosonde-related studies use data from the Integrated Global Radiosonde Archive (IGRA). In contrast to the original high-resolution soundings, IGRA provides a consolidated dataset with a coarser and non-uniform vertical resolution, as it reports values at standard levels (Durre et al., 2006, 2018), and the ERA5 137-level model, however, has a vertical resolution of roughly 300 m at altitudes of 5 km and above, both these radiosonde data and reanalysis/model datasets, while useful for studying global tropopause variations, suffer from limitations in vertical resolutions (Raman and Chen, 2014).

In recent decades, newly developed high vertical resolution (5–10 m) of the radiosonde data allows for fine detailed observation of temperature structure changes within the troposphere and stratosphere. High-resolution sounding data benefits the detection of fine-scale tropopause structures research like multiple tropopause events and provide more reliable estimates of TH changes across hundreds of stations. Furthermore, the global dataset is accumulating, with increasingly long records approaching climate-relevant timescales. Examples include the US (starting in 2005; Ko et al., 2019), China (starting in 2011), and Europe (starting in approximately 1991).

However, reanalysis datasets offer a globally continuous spatial and temporal representation of TH, complementing radiosonde data. Datasets like ERA-Interim are widely recognized for their utility in climate monitoring (Dee et al., 2011), overcoming the spatial resolution limitations of radiosonde data. Therefore, ERA5 has naturally been subjected to comparisons and evaluations with radiosonde in different regions (e.g., Zhu et al., 2021; Velikou et al., 2022; Hoffmann and Spang., 2022). However,

global comparative analyses of high-resolution (10 m) radiosonde data with ERA5 at the tropopause level remain scarce.

Consequently, this study addresses the following questions: (1) How does ERA5 perform in terms of obtaining TH climatologies compared to high-resolution detection? (2) What the global long-term change of TH according radiosonde data and ERA5 reanalysis data? To this end, section 2 details the data and the methods we used. Section 3 presents the comparative analysis and long-term change of TH derived from radiosonde observations and ERA5 reanalysis data. Section 4 ends with a short summary.

2. Data and method

2.1 Radiosonde

Radiosondes are fundamental and crucial data sources for numerical weather prediction models. They are typically carried aloft by weather balloons and burst at altitudes of approximately 27 km (Kumar, 2023). As the radiosonde ascends, it transmits meteorological data including temperature, pressure, relative humidity and air pressure to ground, sea, or air-based receiving stations (Durre et al., 2006). Radiosondes are launched from approximately 800 sites worldwide, regularly twice a day (Ingleby et al., 2016; Durre et al., 2018). Radiosonde data are widely used in studies of planetary boundary layer height (Sorbjan and Balsley, 2008; Seidel et al., 2010), tropopause structure (Birner, 2006; Seidel and Randel, 2006; Añel et al., 2008; Sunilkumar et al., 2017) and gravity waves (Ki and Chun, 2010; Yoo et al., 2018). Key advantages of radiosonde data include in-situ measurements, high vertical resolution, and general reliability. Limitations, however, include relatively low and uneven spatial resolution, as well as a limited temporal frequency of typically twice daily measurements.

Following Guo et al. (2021) and Zhang et al. (2022), we utilized a high-vertical-resolution radiosonde (HVRRS) dataset spanning 2000 to 2023, compiled from multiple sources including the China Meteorological Administration (CMA), the National Oceanic and Atmospheric Administration (NOAA) of the United States, the Deutscher Wetterdienst (Climate Data Center), the Centre for Environmental Data Analysis (CEDA) of the United Kingdom, the Global Climate Observing System (GCOS) Reference Upper-Air Network (GRUAN), and the University of Wyoming.

The original data, sampled at 1-2 second intervals, have a native vertical resolution of approximately 5-10 meters. These data were then uniformly resampled at 10-meter intervals using cubic spline interpolation. This interpolation method provides a good balance between smoothness and accuracy at the scale of interest. Given that the resampling grid spacing (10 m) is comparable to the original data spacing (5-10 m), the potential for spurious oscillations is minimal. Data acquisition was typically performed at 00 UTC and 12 UTC each day. To ensure data quality, only records with at least 10 days (at least one record per day) for which the World Meteorological Organization (WMO) definition of the first tropopause was derived were considered to be valid month and radiosonde stations with at least 10 valid months per year were considered to be valid in that year. Two distinct data selection criteria were established according to study objectives: (1) For radiosonde-ERA5 intercomparison, we analyzed 222

stations (≥ 5 valid years each) with 1,530,517 vertical profiles during years 2000–2023; (2) For long-term TH trend detection, we had 109 stations (≥ 10 valid years each) containing 1,103,730 vertical profiles (Figure 1).

2.2 ERA5

ERA5 is the latest fifth-generation global atmospheric reanalysis product, stands out as one of the best high-resolution atmospheric reanalysis products currently available, utilizing the ECMWF Integrated Forecasting System (IFS) Cy41r2, combined with a 4D-Var assimilation scheme (Hersbach et al., 2020). This remarkable initiative within the Copernicus Climate Change Service (Thépaut et al., 2018) benefits from advancements in modeling and data assimilation over a decade, providing a comprehensive and high-quality record of essential climate variables (Raoult et al., 2017). **Key technical specifications of ERA5 include an hourly temporal resolution, a horizontal resolution of approximately 31 km ($0.25^\circ \times 0.25^\circ$), and 137 hybrid sigma-pressure levels from the surface up to 0.01 hPa (about 80 km height).**

Hoffmann and Spang. (2022) employed the WMO definition to calculate global THs from ERA5 data, making this ERA5-based product available for research purposes. This study utilizes the 2000–2023 ERA5-based data, characterized by a horizontal resolution of $0.3^\circ \times 0.3^\circ$ and a temporal resolution of 1 hour.

2.3 Multivariate regression model for trend analysis

We assessed interannual variability (Section 3.2) using monthly height time series from 2000 to 2023. To investigate the variability driven by tropospheric and stratospheric forcing, the monthly mean TH, $y(t)$, is described by the following multiple linear regression model:

$$y(t) = a \cdot t + B_1 \cdot \text{QBO1} + B_2 \cdot \text{QBO2} + C \cdot \text{ENSO} + D \cdot \text{VOL} + \sum_{k=1}^{11} E_k \cdot \text{Month}_k \quad (1)$$

Here, t is the time in months. The coefficient a represents the linear trend. QBO1 and QBO2 are two orthogonal time series representing the variability of the Quasi-Biennial Oscillation (QBO), constructed from the first two Empirical Orthogonal Functions (EOFs) of stratospheric wind data from the Freie Universität Berlin radiosonde stratospheric winds (Naujokat, 1986). ENSO (El Niño/Southern Oscillation) is represented by the multivariate ENSO Index (source: <https://psl.noaa.gov/enso/mei/>, last access: October 2025). VOL denotes the stratospheric aerosol optical depth at 525 nm, obtained from the Global Space-based Stratospheric Aerosol Climatology project (Thomason et al., 2018). The Month_k are monthly dummy variables ($k = 1, \dots, 11$), which are used to eliminate the seasonal influence on tropopause height variations. B_1 , B_2 , C , D and E_k are the corresponding regression coefficients. Significance is tested based on a two-tailed test with a 95% confidence interval.

2.4 WMO-defined tropopause

The WMO tropopause definition is more robust and generally applicable across a wider range of latitudes. Therefore, calculating TH from radiosonde data using the WMO definition is currently the most suitable method for global tropopause comparisons. As provided by the WMO:

“(a) The first tropopause is defined as the lowest level at which the lapse rate decreases to 2 °C/km or less, provided also the average lapse rate between this level and all higher levels within 2 km does not exceed 2 °C/km;

(b) if above the first tropopause the average lapse rate between any level and all higher levels within 1 km exceed 3 °C/km, then a second tropopause is defined by the same criterion as under (a). This tropopause may be either within or above the 1 km layer.”

The lapse rates are calculated as follows:

$$\Gamma(z_i) = -\frac{\delta T}{\delta z} = -\frac{T_{i+1}-T_{i-1}}{z_{i+1}-z_{i-1}} \quad (2)$$

with T represents temperature, and z represents geopotential height.

Following the WMO definition, Figure 2 presents four examples of radiosonde temperature profiles alongside the corresponding profiles from the nearest ERA5 grid point and its four adjacent points. Cases (a) and (c) show good agreement in tropopause identification. In contrast, cases (b) and (d) reveal significant discrepancies despite having matched overall temperature profiles. In these instances, the high vertical-resolution radiosonde detects a distinct inversion layer—a fine-scale structure not captured by ERA5—which results in a much lower tropopause height (TH). These cases show fine-scale structure could complicate tropopause detection, implying that the existing WMO definition could be further refined for high-resolution dataset.

3. Result

3.1 Radiosonde-ERA5 tropopause height comparison and its seasonal variation

Figure 3 demonstrates that the THs derived from the radiosonde observations and ERA5 reanalysis data are in strong agreement. The linear regression equation ($y = 0.96x + 0.55$) derived from the kernel density scatter plot (Figure 3a) closely follows the 1:1 line. Close agreement between radiosonde and ERA5 is observed for THs around 13 km, with a slight positive bias in radiosonde. A pronounced discrepancy emerges at greater heights (e.g., above 13 km), characterized by increased scatter and deviation from the 1:1 line, forming a distinctive cross-like pattern. Figure 3a reveals several notable outliers, particularly where ERA5 indicates a high tropopause (15–17 km) while the radiosonde data suggest a much lower tropopause height (e.g., 5–6 km). Three potential causes could lead to these outliers. First, the WMO lapse-rate definition is highly sensitive to fine-scale structures in high-resolution data, such as strong inversions, which can be misinterpreted as a tropopause. Second, pronounced inversions associated with phenomena like large-scale subsidence or strong fronts may genuinely satisfy the tropopause criteria at a lower altitude. Third, despite quality control, subtle data issues cannot be entirely

ruled out. This might suggest that both datasets exhibit certain limitations in capturing the TH within the subtropical tropopause break region. Figure 3b displays the distribution of differences between radiosonde and ERA5. On average, TH derived from radiosondes is 32 m higher than that from ERA5, with a mean absolute difference of 336 m. The root mean square error is 756 m, while the Pearson correlation coefficient reaches 0.958 ($p < 0.05$), indicating a statistically significant agreement. Overall, the discrepancy between ERA5 and radiosonde data is minimal, demonstrating strong consistency between the two datasets.

Figure 4 presents the global distribution of mean differences (radiosonde minus ERA5) and mean absolute differences between the two datasets. Across all stations, 85% (187 stations) showed higher mean THs in radiosonde data compared to ERA5. At 206 stations (93%), the mean absolute difference remains within 500 m. Table 1 details the statistical differences for each year from 2000 to 2023, revealing a gradual increase in observation data over time. Intercomparison statistics, detailed in Table 1, reveal that the ERA5-radiosonde TH differences are not homogeneous over time. Instead, the record is marked by two transition points linked to major updates in the ERA5 system. The most pronounced shift occurs around 2006, where the mean difference decreased and the mean absolute difference increased, this change might be due to insufficient GNSS radio occultation data prior to 2006. As a result, ERA5 exhibits a significant cold bias in its stratospheric temperature analysis for the period 2000–2006. To address this, ERA5.1 reanalysis data was specifically generated, employing background error covariance applicable to the 1979–1999 period to correct this bias (Simmons et al, 2020). A secondary, subtler shift is evident around 2016, where the mean difference increased and the mean absolute difference began to decrease, concurrent with the use of the IFS Cy41r2 model cycle as the foundation for ERA5. Changes in the core model physics and assimilation settings provide a plausible explanation for this later discontinuity. Therefore, users of ERA5 TH data should be aware that its homogeneity is influenced by these key evolutionary milestones.

Figure 5 presents 24-year (2000–2023) seasonal mean TH data from radiosonde and ERA5 across different latitudinal bands. Radiosonde and ERA5 data were divided into seven climate zones: Northern Hemisphere/Southern Hemisphere polar (70° – 90°), Northern Hemisphere/Southern Hemisphere mid-latitude (40° – 70°), Northern Hemisphere/Southern Hemisphere subtropics (20° – 40°), and tropics (20° S– 20° N) (Houchi et al., 2010). Seasonal analysis was performed on the station data within each latitudinal band. The uneven and sometimes sparse distribution of radiosonde stations means that their density varies significantly across latitude bands (as indicated on the right side of each latitudinal title in Figure 5). For example, the 40° N– 70° N band contains the most stations (109 stations), while the 70° S– 90° S band has only two stations. This variation makes it challenging to compute a representative regional mean directly from the station data. Although the primary focus of the analysis is the point-to-point ERA5 data, the ERA5 zonal mean (ERA5-F) provides a comprehensive background reference. Except in the tropics and Southern Hemisphere polar region, the seasonal TH variations show consistent patterns: TH reaches maximum values during autumn (June–July–August) and winter (September–October–November), and minimum values during spring (December–January–February) and summer

(March–April–May). This pattern is observed in all regions except the tropics and the Southern Hemisphere polar region.

3.2 Tropopause height long-term trend derived from Radiosonde and ERA5

For long-term trend analysis, we used data with at least 10 years of valid records. Figure 6 presents the annual mean trends over a 24-year period from ERA5 globally. In Figure 6a, the long-term trend of the TH derived from ERA5 data shows a strong latitudinal dependence. Approximately 77% of the regions exhibit a positive trend, with a global average increase (ERA5-F) of +3 m/year. In contrast to the latitudinal distribution of TH—which is higher at low latitudes and lower at high latitudes—the long-term trend generally displays an opposite pattern: lower trends at low latitudes and higher trends at high latitudes. According to ERA5 global data, the Southern Hemisphere exhibits more regions with significant weakening trends than the Northern Hemisphere. Figure 6b compares the long-term trends of the TH between radiosonde and ERA5. The results indicate that the globally averaged TH trend is higher in radiosonde data (+9 m/year) than in ERA5 (+7 m/year), which is consistent with the finding that 83% of individual radiosonde stations also show higher trends.

Figure 7 displays the annual variations of TH across different latitudinal bands from 2000 to 2023, as observed by radiosonde and ERA5. The results show strong agreement between radiosonde and ERA5 in terms of the annual mean TH. The figure further shows that the coverage periods of high-resolution radiosonde data deemed suitable for long-term trend analysis according to Section 2.1 vary across different latitude zones, with the tropical zone (20°S–20°N) having the shortest record (2010–2023). Multiple regression analysis (Table 2) reveals a significant global-scale increase in TH and distinct latitudinal patterns in the annual trends. The most robust increase (+15 m/year for radiosonde and +13 for ERA5) occurs in the Northern Hemisphere polar (70°N–90°N), while the Southern Hemisphere polar (70°S–90°S) exhibits a unique and sharp decline (-13 m/year for radiosonde and -11 m/year for ERA5). Serving as a spatial background, the ERA5-F is characterized by a stronger upward trend in the Northern Hemisphere, while the Southern Hemisphere exhibits an insignificant upward or downward trend. The long-term trend values for TH derived from radiosonde observations and spatially/temporally collocated ERA5 data show high consistency across latitudinal bands. However, due to the limited number of stations, the trends in the following regions should be interpreted with caution: 70°S–90°S, 40°S–70°S, and 70°N–90°N.

4. Conclusion and discussion

TH is an indispensable metric in climate change research, directly influenced by the temperature structure of the troposphere and stratosphere. Direct observations of the tropopause via radiosonde are generally most reliable, but are limited by the uneven global distribution of measurement stations, posing a significant constraint on global climate change studies. Recent reanalysis studies (e.g., Xian and Homeyer, 2019; Tegtmeier et al., 2020; Hoffmann and Spang, 2022; Zou et al., 2023) have advanced our understanding of TH variability, systematic evaluations of their biases and spatiotemporal patterns

285 against high-resolution radiosonde data have yet to be fully conducted, particularly for newer reanalyses
such as ERA5. While radiosondes represent the most reliable data source, they suffer from stringent
spatiotemporal limitations. In contrast, the ERA5 reanalysis dataset provides complete spatiotemporal
coverage and has been widely utilized in atmospheric studies. Therefore, a systematic intercomparison
between high-resolution radiosonde and ERA5 THs is essential to both reconcile dataset discrepancies
290 and optimize their combined use for robust trend analysis.

This study investigates the spatial and temporal discrepancies between ERA5-derived THs and
high-resolution radiosonde data. Building upon previous research by Xian and Homeyer (2019) and
Hoffmann and Spang (2022), the present study confirms an overestimation (32 m) of TH in radiosonde
compared to ERA5, a bias that falls within the range of uncertainties reported in earlier intercomparison
295 efforts. Furthermore, we highlight that the WMO-defined thermodynamic tropopause may share
mathematical similarities with thin, low-altitude temperature inversions detected in high-resolution
radiosonde profiles, introducing complexity in TH determination. In the statistical analysis, the
differences between radiosonde and ERA5 are more pronounced in the subtropical region. This
discrepancy may be linked to the subtropical tropopause break phenomenon, suggesting a region-specific
300 bias that warrants attention in future reanalysis developments. Regarding the seasonal variation of TH,
except in the tropics and the Southern Hemisphere polar region, TH peaks in autumn and winter while
being lower in spring and summer (seasons are reversed between hemispheres). However, radiosonde
and ERA5 exhibit strong consistency in TH, with only minor differences about 32 m.

Comprehensive validation against high-resolution radiosonde observations demonstrates ERA5's
305 exceptional performance in capturing TH characteristics, including absolute values, temporal variations,
and spatial correlations. Application of the WMO tropopause definition to high-resolution soundings
reveals potential limitations, particularly in subtropical regions where strong temperature inversions may
lead to detection artifacts. These thin but intense inversion layers are often unresolved by coarse-
resolution temperature profiles.

310 The global tropopause elevation trend is quantified at +90 m/decade based on radiosonde
measurements. Given the 10-m vertical resolution of the sounding data, this rising trend represents a
statistically robust signal, while ERA5 shows a trend of +70 m/decade, with 83% of radiosonde stations
exhibiting higher trends than ERA5. Point-to-point comparisons indicate ERA5 systematically
underestimates long-term TH trends by approximately 40% relative to radiosonde benchmarks. This
315 discrepancy warrants further investigation through coordinated model-observation intercomparison
studies. The ERA5-F global analysis reveals that 77% of regions exhibit an increasing trend in TH. The
global TH exhibits a significant but spatially uneven upward trend, with hemispheric asymmetry. The
Arctic region shows the most pronounced increase (+15 m/year for radiosonde and +13 for ERA5), while
the Antarctic exhibits a sharp downward trend (-13 m/year for radiosonde and -11 for ERA5).

320 However, radiosonde and ERA5 data demonstrate high consistency, yet limited station coverage
(particularly in the Southern Hemisphere) remains a primary source of uncertainty, necessitating cautious
interpretation of regional trends. The variation of TH with climate exhibits significant latitudinal and

longitudinal differences. But many regions lack high-resolution, continuous radiosonde observations for cross-validation. This necessitates further verification through expanded station coverage and longer-term observations, **which are identified in this and other studies as critical for understanding TH dynamics.**

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Competing interests

The contact author has declared that neither they nor their co-authors have any competing interests.

Data availability

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Author contributions

JZ conceptualized this study. YG carried out the analysis with comments from other co-authors. YG and JZ wrote the original manuscript. WW, SZ provided useful suggestions for the study. All authors contributed to the improvement of paper.

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Table 1. Statistical analysis by years, RS-E5: Radiosonde minus ERA5-based, Bias: average of differences (ERA5-based minus Radiosonde), Abs-bias: average of absolute values of differences, RMSE: root mean square error, R*: Pearson correlation coefficient ($p < 0.05$).

Year	Valid record	Bias(RS-E5) (m)	Abs-bias(RS-E5) (m)	RMSE (m)	R*
2000	4651	44	243	460	0.938
2001	7120	43	236	446	0.957
2002	7335	56	228	376	0.969
2003	7377	40	231	417	0.962
2004	7739	48	227	401	0.964
2005	8105	35	267	527	0.943
2006	11259	35	302	632	0.960
2007	25650	14	358	781	0.951
2008	40154	20	362	795	0.953
2009	47716	-28	422	939	0.929
2010	54066	20	347	764	0.960
2011	72234	-12	399	918	0.938
2012	71474	-20	387	906	0.941
2013	71930	-1	375	885	0.945
2014	78010	22	366	837	0.952
2015	87345	28	350	771	0.959
2016	80738	21	381	861	0.952
2017	89145	51	336	732	0.962
2018	153446	46	312	685	0.966
2019	126090	57	286	621	0.972
2020	120414	58	285	612	0.972
2021	103729	45	375	852	0.948
2022	110719	51	304	662	0.966
2023	144071	47	304	692	0.964
Total	1530517	32	336	756	0.958

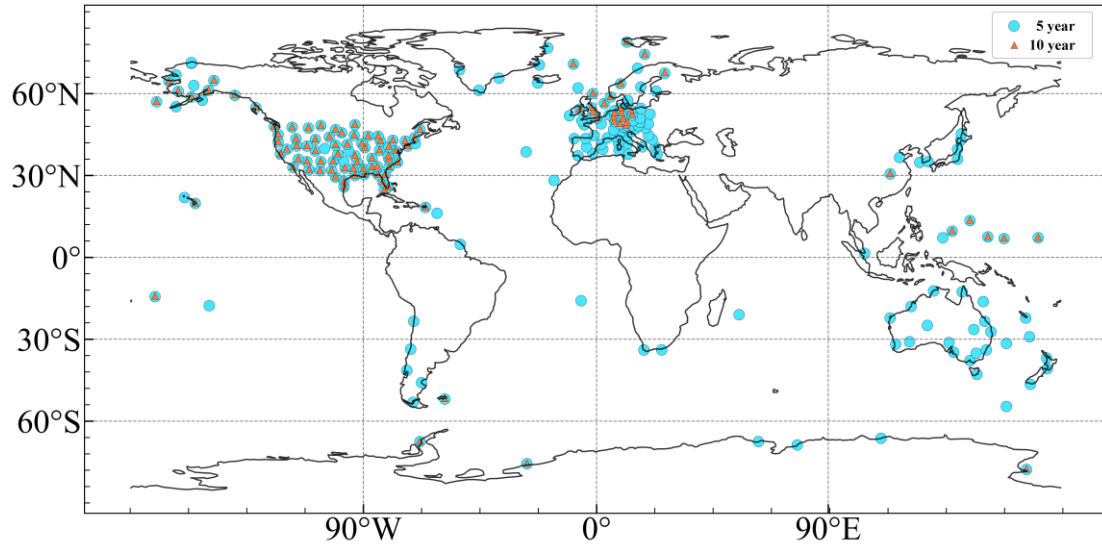
510 **Table 2.** Tropopause height trend from different latitude zones (stations with ≥ 10 years valid records; 40 °S–70 °
S excluded due to data absence). ERA5 is a one-dimensional data that corresponds one-to-one in time and space
with the Radiosonde data, and ERA5-F is the overall mean of the ERA5 global trend averages over the period
2000 to 2023 in a two-dimensional plane divided according to latitudinal bands.

Latitudinal zone	Number of stations	Radiosonde (m/year)	ERA5 (m/year)	ERA5-F (m/year)
90 °N–70 °N	3	15	13	10
40 °N–70 °N	56	8	7	6
20 °N–40 °N	37	11	8	4
20 °S–20 °N	8	11	7	3
20 °S–40 °S	0	\	\	-2
40 °S–70 °S	2	5	7	-1
70 °S–90 °S	2	-13	-11	2
Global	108	9	7	3

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Figure 1. Spatial distribution of the qualifying radiosonde stations used in this study. The number of qualifying radiosonde stations with valid data records for 10 years (indicated by orange triangles) is 108, and that for 5 years (indicated by cyan dots) is 222. The United States and Europe are data-intensive.

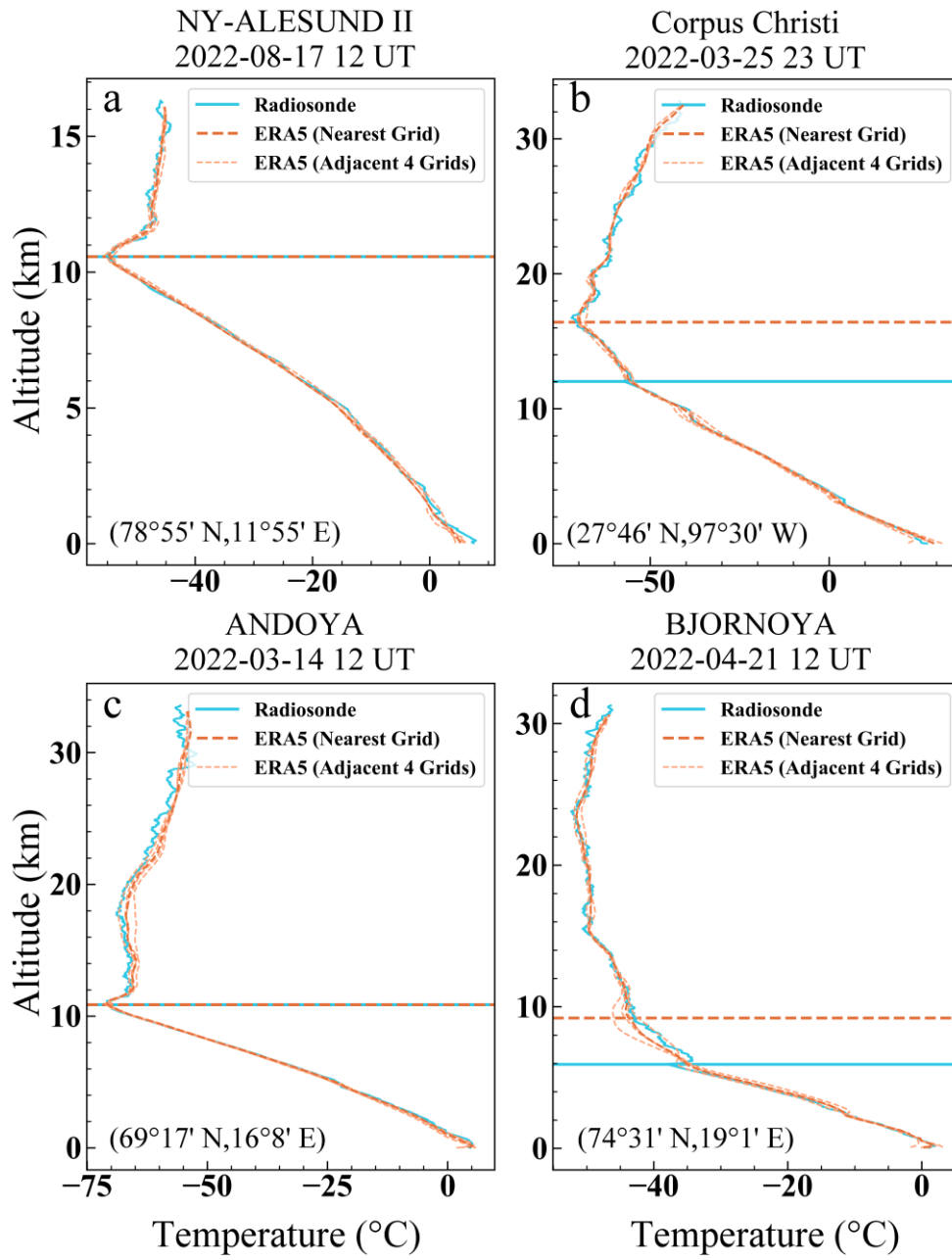


Figure 2. Tropopause height identification from Radiosonde and ERA5 temperature profiles. The Radiosonde profile (cyan) is compared against the nearest ERA5 grid point (solid orange) and its four adjacent points (thin orange). The diagnosed tropopause is marked by a horizontal line. The Radiosonde location, date, and time are indicated in each subplot.

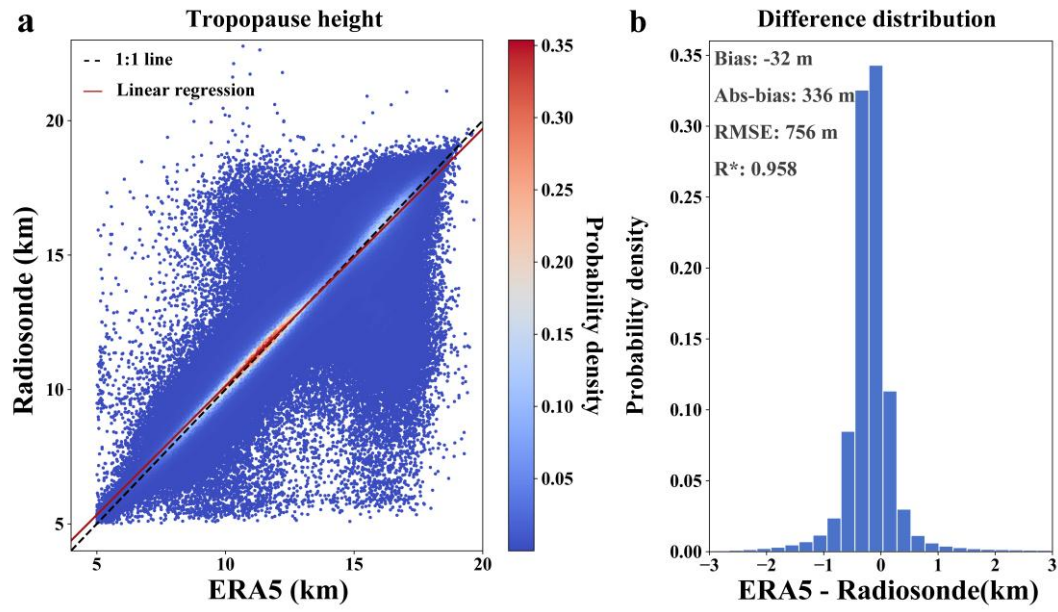


Figure 3. Kernel density scatter plot (a) and probability density histogram (b) comparing tropopause heights. The black dotted line indicates a 1:1 relationship, the red line shows the linear regression ($y = 0.96x + 0.55$). Bias (average of differences, ERA5 minus Radiosonde), absolute bias (average of absolute values of differences), RMSE (root mean square error), and R (Pearson correlation coefficient, $p < 0.05$) are also indicated.

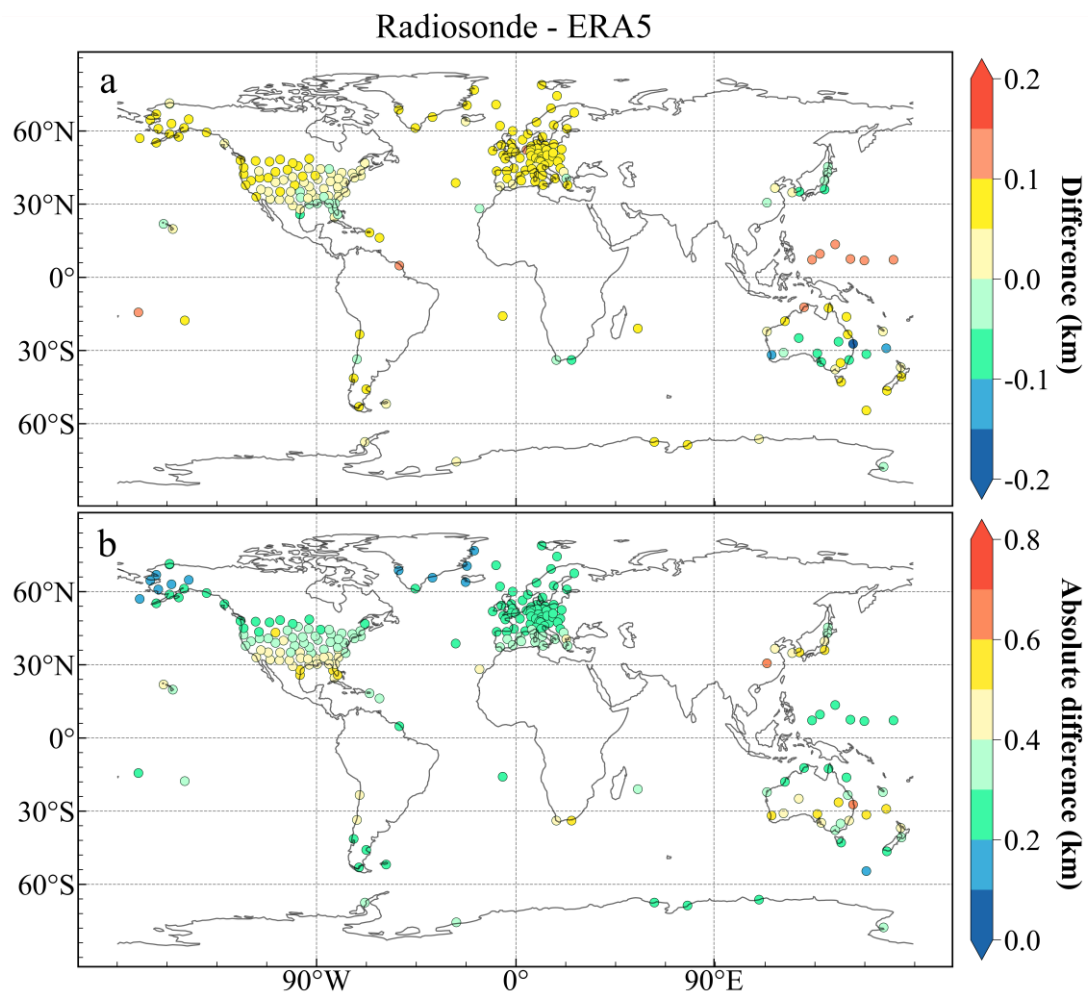


Figure 4. Global distribution of station-mean differences between radiosonde observations and ERA5 (a) and station-mean absolute differences (b).

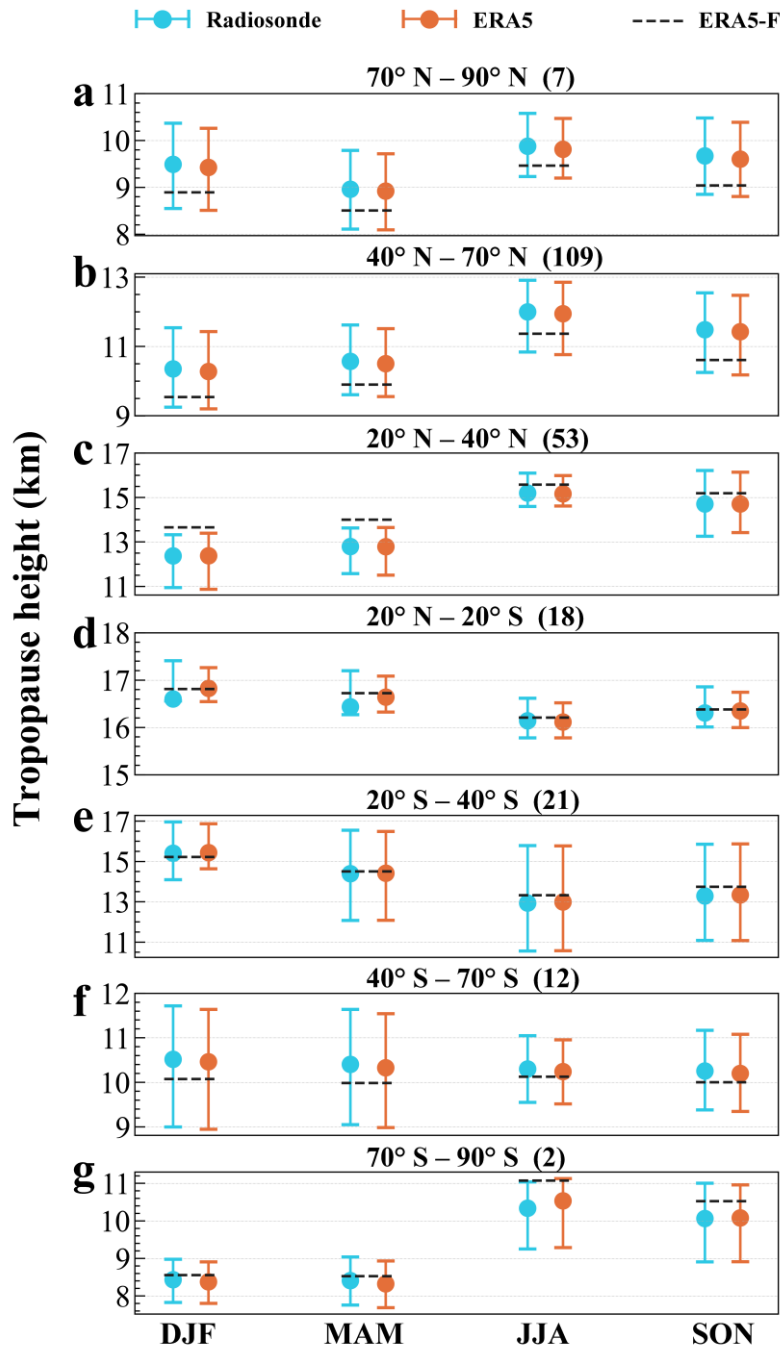


Figure 5. Seasonal change across 7 latitudinal bands. The number of stations on the right side of the latitude band titles. ERA5 is a one-dimensional data that corresponds one-to-one in time and space with the radiosonde data, and ERA5-F is the overall mean of the ERA5 global seasonal averages over the period 2000 to 2023 in a two-dimensional plane divided according to latitudinal bands. DJF: December–January–February, MAM: March–April–May, JJA: June–July–August, SON: September–October–November. The dots represent the mean, and the bar represents the 25% or 75% quantiles.

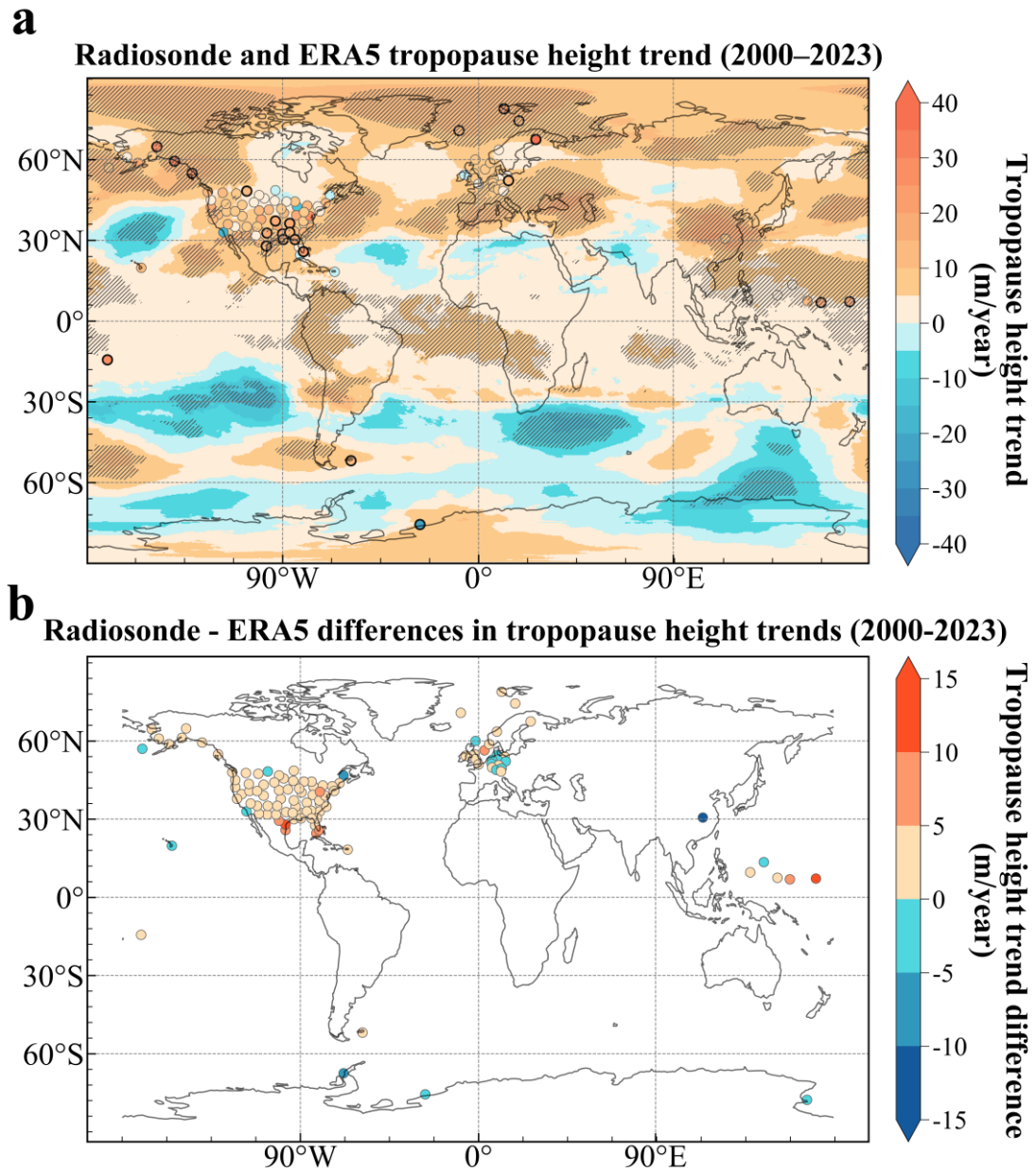


Figure 6. Global map of tropopause height trend (a) and its difference (b) using radiosonde and ERA5 data. The shaded area (ERA5) and black thick ring (radiosonde) in panel (a) indicate statistically significant trends in tropopause height ($p < 0.05$), based on the multiple regression analysis at each ERA5 grid point and radiosonde station.

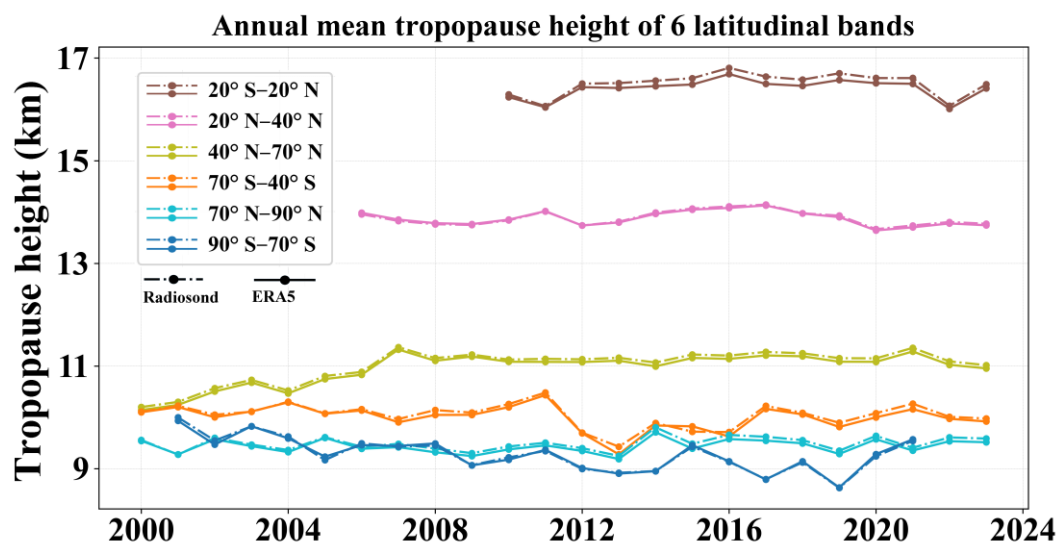


Figure 7. The time series are shown for six latitude bands (no data available for 20 °S–40 °S). Data from radiosonde and ERA5 are depicted by dashed and solid lines, respectively.