

The manuscript entitled “Can high-resolution convection-permitting climate models improve flood simulation in southern Quebec watersheds?” investigates whether CPRCM simulations provide added value for flood simulation in southern Quebec, with focus on extreme summer and fall rainfall events. The topic is relevant and the objectives of the study are clear. The manuscript is also well organized and written. The comparison between the 12 km and 2.5 km CRCM6/GEM5 simulations, and try to link precipitation extremes to hydrological responses, are interesting to both the regional climate and hydrological communities. However, several important issues need to be addressed before the conclusions can be fully supported.

1. The study aims to assess the added value of CPRCM simulations. In this context, the choice of a lumped hydrological model raises concerns. The hydrological model used in this study (GR5dt) is a conceptual and lumped model with only elevation bands. Given that the main added value of CPRCMs lies in their improved representation of the spatial patterns of intense precipitation, the use of a lumped model may substantially limit the ability to assess this added value. For example, Figure 5 shows that for many basins the KGE driven by RCM-12 km is higher than that driven by CPRCM-2.5 km, which may partly reflect this modeling choice. While this limitation is mentioned in the discussion, the authors should better justify the choice of model and more clearly acknowledge the associated limitations when interpreting the improvements in flood simulations.

We thank the reviewer for this important comment. We agree that the use of a conceptual lumped hydrological model such as GR5dt limits the ability to fully exploit the spatial information provided by CPRCM simulations.

The choice of GR5dt was motivated by the objectives of this study, which focus on assessing the sensitivity of simulated flood responses to differences in precipitation intensity and temporal structure between CPRCM (2.5 km) and coarser-resolution RCM (12 km) outputs, rather than on explicitly resolving spatial runoff generation processes within catchments. Importantly, only a limited number of hydrological models are suitable for simulations at the hourly time scale, and GR5dt is one of the few lumped models specifically designed and validated for this purpose. GR models have been widely used and shown to perform robustly for flood simulations at daily and sub-daily time scales across a range of catchment sizes, while requiring limited calibration parameters, which was essential for ensuring a consistent comparison across multiple basins.

We fully acknowledge that lumped models cannot explicitly represent the spatial variability of rainfall and runoff processes within a basin. However, the catchments considered in this study are generally small and characterized by relatively flat topography, conditions under which lumped modeling approaches are commonly considered appropriate.

Nevertheless, part of the added value of CPRCMs, particularly the improved representation of fine-scale spatial precipitation patterns, may not be fully translated into improved simulated river discharge when using a lumped framework. This limitation may partly explain why, for several basins, the KGE driven by the 12 km RCM exceeds that obtained using the 2.5 km CPRCM (Fig. 5).

To address this point more explicitly, lines 568-571 will be replaced by following paragraph:

'The use of a lumped hydrological model in this study was a deliberate and conservative choice aimed at maintaining the methodology simple and enhancing the reproducibility. By design, lumped models do not explicitly take advantage of the spatial organization of precipitation, and therefore the fine-scale rainfall structures simulated by convection-permitting climate models (CPRCMs) are not directly translated into the runoff generation process. Consequently, the hydrological response primarily reflects differences in precipitation intensity and temporal variability, while spatially localized convective extremes are averaged at the catchment scale when used in a lumped model. This aggregation may attenuate or even mask potential CPRCM added value, and likely contributes to cases where the 12 km simulation yields higher KGE values than the 2.5 km simulation. Moreover, the fact that hydrological improvements are nevertheless observed using CPRCM outputs within a lumped modeling framework indicates that these results represent a lower-bound estimate of the added value of high-resolution climate simulations. This suggests that greater gains in the simulation of extreme flows could be achieved in future studies through the use of distributed hydrological models, such as Hydrotel (Fortin et al., 2001), which are better suited to exploit spatial rainfall heterogeneity. Overall, these findings underscore that hydrological added value depends not only on atmospheric model resolution, but also on the consistency between climate and hydrological model structures, and that the limitations of the lumped approach should be explicitly considered when interpreting improvements in flood simulations.'

In addition, the manuscript states that "KGE values slightly increase as basin size

increases”, whereas the figure shows substantial variability from left to right. Other factors may influence the results, such as elevation-related precipitation biases. An elevation-dependent analysis (e.g., bias or performance metrics versus basin mean elevation) could be very informative, especially given known elevation-dependent biases in precipitation.

We thank the reviewer for this observation. We agree that the statement suggesting that KGE values increase with basin size was too strong given the substantial variability observed across basins. Our intention was to indicate a weak tendency rather than a systematic relationship. To avoid overinterpretation, we will revise the text to clarify that basin size alone does not explain the variability in model performance.

We further agree that additional factors, such as elevation-related precipitation biases, may influence the hydrological results. Orographic effects and elevation-dependent biases are known to affect both precipitation magnitude and phase, and their impact may differ between the 2.5 km and 12 km simulations. In the current study, these effects are not explicitly analyzed, and we therefore interpret the hydrological results as reflecting an interplay between basin characteristics, precipitation forcing, and model structure rather than a simple dependence on basin size.

We will revise the Discussion to acknowledge the potential role of elevation and to highlight that an elevation-dependent evaluation (e.g., relating bias or performance metrics to basin mean elevation) would provide additional insight into the hydrological added value of CPRCM simulations. Such analyses are identified as an important direction for future work.

2. Table 3 presents useful results. However, in Figure 7, several basins do not show a clear added value in the boxplots of peak flow bias (e.g., small basins such as Bras d’Henri, as well as larger basins such as Eaton, Au Saumon, and Etchemin). Additional explanation would be helpful here. In addition, it would be useful to clarify whether the peak flow bias reported in Table 3 is defined consistently with that shown in Figure 9.

We thank the reviewer for this careful reading of the results. We agree that the added value of the CPRCM-driven simulations is not systematic across all basins, as illustrated by the boxplots of peak flow bias in Figure 7. This heterogeneity reflects differences in basin characteristics, scale, and the interaction between precipitation forcing and the hydrological modeling framework.

For small basins such as Bras d'Henri, flood response is highly sensitive to localized precipitation extremes and their spatial positioning within the catchment. While the CPRCM provides a more realistic representation of localized convective precipitation, this added spatial detail is largely averaged out when used as input to the lumped hydrological model, which may limit or even counteract potential improvements in peak flow simulation.

For larger basins (e.g., Eaton, Au Saumon, and Etchemin), the basin-averaging effect becomes even stronger, and flood peaks are influenced by a combination of spatial aggregation, routing processes, and also on how runoff from different parts of the basin reaches the outlet at similar or different times i.e. temporal synchronization of runoff contributions. In such cases, smoother precipitation fields from the coarser-resolution RCM may sometimes lead to comparable or better peak flow bias metrics, despite a less realistic representation of precipitation extremes at smaller scales.

Yes, the peak flow bias reported in Table 3 is calculated using the same peak flow values as those shown in Figure 9. To avoid any ambiguity, we will clarify this definition in the manuscript.

3. Biases in simulated precipitation strongly affect hydrological model performance. I understand that no bias correction was applied to the climate model forcings before driving the hydrological model. However, the potential impact of precipitation biases on the flood simulation results, should be discussed more explicitly.

We thank the reviewer for this important comment. We agree that biases in simulated precipitation can strongly influence hydrological model performance, particularly for flood simulations. In this study, no explicit bias correction was applied to the climate model forcings, as the objective was to assess the relative hydrological response to precipitation generated by climate models at different spatial resolutions within a consistent modeling framework. Bias correction would add another layer of complexity that might make the results difficult to assess. Furthermore, bias correcting subdaily precipitation data at a 2.5 km resolution is not a trivial matter, especially with regards to correcting extremes as well as the rarity and uncertainty related to the reference dataset that would need to be used.

We acknowledge that biases in precipitation magnitude, intensity, and temporal structure may directly affect simulated flood peaks and performance metrics. In particular, biases in short-duration extreme precipitation may lead to over- or

underestimation of peak flows, depending on basin characteristics and hydrological model sensitivity. While both the 2.5 km and 12 km simulations are affected by model biases, their impacts may differ across basins and seasons, contributing to the heterogeneous hydrological performance observed in the results.

To address this point more explicitly, we will expand the Discussion to clarify how precipitation biases may influence the flood simulation results and to emphasize that the hydrological outcomes should be interpreted primarily in a comparative sense rather than as bias-free representations of observed floods. The application of bias-correction techniques and their interaction with model resolution are identified as important directions for future work.

4. The manuscript suggests that CPRCM outputs improve flood simulations and can be useful for risk management strategies. It remains unclear whether the authors imply that CPRCM outputs can be used directly for hydrological applications. If so, under what conditions (e.g., bias correction, calibration)? A clearer discussion of the applicability and limitations of using CPRCM outputs for hydrological modeling would strengthen the paper.

We thank the reviewer for this important comment. We clarify that we do not argue that CPRCM outputs are universally bias-free; however, our results indicate that CPRCM simulations show closer agreement with station observations than the coarser-resolution RCM, particularly for extreme precipitation. In this context, we believe that CPRCM outputs can be used directly for flood simulations to provide a more realistic estimation of flood magnitudes, especially in comparative or exploratory studies. Bias adjustment of the bulk of the distribution is generally more straightforward than that of the extremes. If, for example, the biases in precipitation extremes simulated by the CPRCM are shown to be acceptable, it may be justified to use the data within a hybrid correction framework in which the bulk of the distribution is bias-corrected (to preserve the annual water balance), while the extremes are left unadjusted. This highlights the importance of improving our understanding of biases in extreme precipitation. However, the development of such a correction approach is beyond the scope of the present study.

At the same time, we acknowledge that systematic precipitation biases may still be present and that bias correction can further improve hydrological performance, particularly for applications requiring accurate absolute flood quantification or in cases where the biases are too large to ignore.

Bias correction is therefore recommended for operational or risk management applications, but its application is not strictly required to assess relative flood behavior or to benefit from the improved representation of precipitation extremes provided by CPRCMs.

We will revise the Discussion to clarify the conditions under which CPRCM outputs may be used directly and to more clearly distinguish between exploratory flood analyses and applications requiring high accuracy, such as flood risk management.

Overall, the study addresses an important question, and provides useful information for hydroclimate research.