https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

Measurement of soluble aerosol trace elements: inter-laboratory comparison of eight leaching protocols

Mingjin Tang,[1,10,#,*] Morgane M.G. Perron,[2,#] Alex R. Baker,[3,#] Rui Li,[1,#] Andrew R. Bowie,[4,8,] Clifton S. Buck,[5,] Ashwini Kumar,[6,] Rachel Shelley,[3,] Simon J. Ussher,[7,] Rob Clough,[7] Scott Meyerink,[8] Prema P. Panda[6], Ashley T. Townsend,[9] Neil Wyatt[7]

[1] State Key Laboratory of Advanced Environmental Technology and Guangdong Key Laboratory of Environmental Protection and Resources Utilization, Guangzhou Institute of Geochemistry, Chinese Academy of Sciences, Guangzhou, China

[2] UMR 6539 CNRS/UBO/IRD/Ifremer, Laboratoire des sciences de l'environnement marin, Université de Brest – Institut Universitaire Européen de la Mer, 29280 Plouzané, France

[3] Centre for Ocean and Atmospheric Science, School of Environmental Sciences, University of East Anglia Norwich, UK

[4] Institute for Marine and Antarctic Studies, University of Tasmania, Tasmania, Australia

[5] Skidaway Institute of Oceanography, University of Georgia, Savannah, GA, USA

[6] CSIR – National Institute of Oceanography, Dona Paula, Goa, India

[7] School of Geography, Earth and Environmental Sciences, University of Plymouth, Plymouth, UK

[8] Australian Antarctic Program Partnership (AAPP), University of Tasmania, Battery Point, Tasmania, Australia

[9] Central Science Laboratory, University of Tasmania, Hobart, Tasmania, Australia

[10] Institute of Surface-Earth System Science, School of Earth System Science, Tianjin University, Tianjin, China

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

# Equal contribution.

* Correspondence. Mingjin Tang (mingjintang@126.com)

Abstract

A range of leaching protocols have been used to measure the soluble fraction of aerosol trace elements worldwide, and therefore these measurements may not be directly comparable. This work presents the first large-scale international laboratory intercomparison study for aerosol trace element leaching protocols. Eight widely-used protocols are compared using 33 samples that were subdivided and distributed to all participants. Protocols used ultrapure water, ammonium acetate, or acetic acid (the so-called "Berger leach") as leaching solutions, although none of the protocols were identical to any other. The ultrapure water leach resulted in significantly lower soluble fractions, when compared to the ammonium acetate leach or the Berger leach. For Al, Cu, Fe and Mn, the ammonium acetate leach resulted in significantly lower soluble fractions than those obtained with the Berger leach, suggesting that categorizing these two methods together as "strong leach" in global databases is potentially misleading. Among the ultrapure water leaching methods, major differences seemed related to specific protocol features rather than the use of a batch or a flow-through technique. Differences in trace element solubilization among leach solutions were apparent for aerosols with different sources or transport histories, and further studies of this type are recommended on aerosols from other regions. We encourage the development of "best practices" guidance on analytical protocols, data treatment and data validation in order to reduce the variability in soluble aerosol trace element data reported. These developments will improve understanding of the impact of atmospheric deposition on ocean ecosystems and climate.

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

1 Introduction

Atmospheric deposition has been a major pathway for vital nutrients, including trace elements, to reach the surface ocean over modern and geological times (Jickells et al., 2005; Jaccard et al., 2013; Mahowald et al., 2018). Natural ocean fertilization events have been reported following aeolian deposition of the vital micronutrient iron (Fe) from dust (Cassar et al., 2007), volcanic emissions (Langmann et al., 2010) and fire emissions (Tang et al., 2021). Since the Industrial Revolution, increasing atmospheric emissions linked to human activities as well as changes in anthropogenic land use have resulted in additional inputs of trace elements into the atmosphere (Mahowald et al., 2018; Bai et al., 2023). Notably, the increased emission of toxic metals (e.g., Cu) has been shown to have the potential to negatively impact marine ecosystems (Paytan et al., 2009; Jordi et al., 2012).

While anthropogenic activities may result in the emission of trace elements which are deleterious to the marine ecosystem, they are also rich in bioavailable essential nutrients, such as Fe (Hamilton et al., 2020; Ito et al., 2021) and Mn (Lu et al., 2024). In addition, atmospheric mixing with anthropogenic pollutants can enhance the solubility of natural aerosol Fe (and perhaps other trace elements) due to proton-promoted and ligand-mediated interactions (Shi et al., 2012; Paris and Desboeufs, 2013; Baker et al., 2021).

The biogeochemical impacts of aerosol trace elements deposited to the ocean are primarily driven by the fraction that is assimilated by the marine microbial community (Baker and Croot, 2010; Jickells et al., 2016; Mahowald et al., 2018). This fraction has historically been related to operational definitions of trace elements released into solution in experimental studies that have employed a wide variety of leaching protocols (e.g., Sholkovitz et al., 2012; Fishwick et al., 2014; Perron et al., 2020a; Li et al., 2023) and have variously described the released trace element fraction as “soluble”, “labile”, “leachable”, “dissolved”, “readily-accessible”, “bioaccessible” and “bioavailable”. Thus, while a considerable number of studies have

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

investigated aerosol soluble trace elements (e.g., Hsu et al., 2005; Baker and Jickells, 2006; Buck et al., 2010; Kumar et al., 2010; Sholkovitz et al., 2012; Gao et al., 2020; Perron et al., 2020b; Chen et al., 2024), data reported in the literature suffers from a lack of standardization of protocols and terminology (Meskhidze et al., 2019). In addition, such operationally defined fractions do not map onto the oceanic definitions of "soluble" or "dissolved" trace elements in seawater, and understanding of the relationships between these fractions and the metabolic processes of nutrient uptake by marine organisms is currently lacking.

Further complication arises due to the solubilization of aerosol trace elements after deposition into the ocean, which can be influenced by varying properties of seawater, as well as through dissolution kinetics during the lifetime of particles in the seawater column that are difficult to replicate in the laboratory (Baker and Croot, 2010). Different leaching protocols may therefore simulate the solubilization of aerosol trace elements under different atmospheric and oceanic conditions as well as over different timescales, but the environmental relevance of the various leaching protocols in use is currently unclear.

The GEOTRACES community has made significant advances in producing a series of recommendations for aerosol sampling, sample handling and sample digestion for total trace element determination (Morton et al., 2013; Aguilar-Islas et al., 2024; Buck et al., 2024). Similar standardization has not been applied to aerosol soluble trace element determination, rendering both comparisons between different studies and data conglomeration for use in modeling studies very difficult (Perron et al., 2024; Shelley et al., 2024).

Several studies (Chen et al., 2006; Mackey et al., 2015; Clough et al., 2019; Perron et al., 2020a; Li et al., 2023) have compared aerosol trace element leaching methods and assessed the effects of different leaching solutions and contact time, although in each of these studies laboratory handling was performed by a single group and analysis was undertaken typically using a single instrument. A recent study (Li et al., 2024) found good agreement between

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

aerosol solubility for eight trace elements measured by two laboratories using four different ultrapure water (UPW) batch leaching methods. Moreover, Li et al. (2024) suggested that very small and sometimes insignificant differences are introduced by varying agitation method, filter pore size and contact time within UPW batch leaching protocols.

Here we present the first large-scale intercomparison study which compares eight commonly used leaching protocols for determining soluble trace elements in aerosol samples. Six research institutions participated in this intercomparison: Guangzhou Institute of Geochemistry (GIG), China; the CSIR-National Institute of Oceanography (NIO), India; the University of East Anglia (UEA), UK; the University of Georgia (UGA), USA; the University of Plymouth (UoP), UK; and the University of Tasmania (UTAS), Australia. Leaching protocols examined use UPW, ammonium acetate (AmmAc) or acetic acid with hydroxylamine hydrochloride (Berger) as leaching solutions; they also varied in contact time, agitation method and filtration procedure. No attempt was made to standardize the sample processing, analysis and the data treatment, with each group using their usual practices for these procedures. The aim of this study is to assess the extent to which the variability in the reported soluble fractions of aerosol trace elements can be attributed to the leaching methodology used and/or sample characteristics.

2 Methodology

2.1 Sample collection

This work used a custom-made instrument (ASM-1) (Jiang et al., 2022; Li et al., 2024) to collect ambient PM_{10} samples with a sampling flow rate of 1 m^3/min. Whatman 41 cellulose fiber filters (203 mm × 254 mm) were used to collect aerosol particles due to their low background (blank concentrations) for trace elements following the recommended GEOTRACES cleaning protocol. The actual area available for aerosol collection was 180 mm × 230 mm due to filter edges being covered by the frame of the filter holder. Filters used for

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

aerosol sampling were acid-washed using procedures described in previous work (Morton et al., 2013; Zhang et al., 2022) to further reduce the background, and stored individually in zipper-top polyethene bags (Sigma-Aldrich, 9 inch × 12 inch).

Six PM_{10} samples were collected at an urban site (113°36'E, 23°13'N) in Guangzhou from 24 November to 01 December 2021. The sampling site is located on the rooftop of a building at the Guangzhou Institute of Geochemistry, about 30 m above ground level (Yu et al., 2020). Sampling started at 08:30 or 20:30 each day (UTC+8), and lasted for 11.5, 23.5 or 35.5 hours to vary the amount of aerosol particles collected. In addition, three lab blanks (i.e., acid-washed filters which were not taken to the field) and four field blanks (i.e., acid-washed filters mounted in the aerosol sampler for 2 h while the pump was off) were prepared during this sampling period.

Another 33 PM_{10} samples were collected between 03 April and 07 May 2022 at a suburban site in Qingdao, a coastal city in Northern China affected by Asian desert dust and anthropogenic pollution. The sampling site is located on the rooftop of a building (36.34°N, 120.67°E), about 20 m above ground level (Zhang et al., 2022). Aerosol sampling started at 08:00 each day and ended at 07:30 on the next day (UTC+8), giving a sampling time of 23.5 h and a sampled air volume of 1410 m^3. In addition, six lab blanks and seven field blanks were collected during the sampling period in Qingdao with the same procedure as described above.

After sampling, each filter was folded inward to protect aerosol particles, placed back into the zipper-top polyethene bag, and stored frozen at -20 °C.

2.2 Sample use and distribution

All the PM_{10} samples and blank filters were divided into eight discs (47 mm in diameter) using a circular titanium hole-punch. These subsamples were folded inward, stored individually in zipper bags (Sigma-Aldrich, 2.5 inch × 3 inch), and labelled X-1 to X-8 (where X is the sample identification name). Subsamples were labelled after subdivision, hence the subsample

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

numbers are not related to the location of the discs on the filter.

Intercomparison exercises should ideally be conducted by sharing homogeneous materials among the participants in quantities that are sufficient for each participant to make a realistic assessment of the precision of their measurements. Neither of these conditions could be easily met for the intercomparison study reported here because: 1) it could not be guaranteed that aerosol material was homogeneously distributed over the aerosol filters; and 2) the amount of material a subsample contained was unlikely to be sufficient for replicate analysis of the soluble trace elements studied. Our study therefore addressed the issues of sample homogeneity (Section 2.6.1), intra-method precision (Section 2.6.2) and inter-method comparison (Section 2.6.3) separately.

For the homogeneity test, all eight portions of each filter collected at Guangzhou (n = 6) were subjected to the same digestion procedure by a single group at GIG (Section 2.3) to determine the total concentration of 14 trace elements in each subsample. This information directly informs whether sample heterogeneity represents a confounding factor in later comparisons.

The 33 samples (and six lab blanks and seven field blanks) collected in Qingdao were distributed to all participating laboratories for soluble trace element analysis. Seven of these samples (C1-C7) were used to assess intra-method precision with each laboratory receiving three subsamples of two filter samples (Table 1). Furthermore, each laboratory received one portion of each of the remaining 26 filter samples (D1-D26) for conducting the leaching method intercomparison. One subsample of each of the Qingdao samples was also retained at GIG for total trace element determination, and the last subsample was reserved for future as yet undetermined usage.

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

Table 1. Summary of subsample distribution from filter samples in group C to the six laboratories. Each lab was provided with triplicates of each C filter.

GIG	UTAS	UGA	UEA	UoP	NIO
C1	C1	C3	C4	C5	C6
C7	C2	C4	C5	C6	C7

2.3 Leaching and digestion procedures

Table 2 provides an overview of the laboratory leaching protocols investigated in this study; more details can be found in Tables S1-S6. UGA and UTAS both employed a two-stage protocol with successive leaching (termed leach 1 and leach 2, hereafter) and therefore reported two values for each trace element. A total of eight leaching protocols were examined in this study.

The eight leaching protocols fell into three broad categories based on the chemistry of the leaching solution: five methods used ultrapure water (> 18.2 MΩ.cm) as leachate (two employing batch extraction and three employing flow-through extraction), while the other three methods used ammonium acetate (two methods) or acetic acid and hydroxylamine hydrochloride (one method) as leachate. These categories are referred to hereafter as UPW, AmmAc and Berger leaches, respectively. Within and among these categories, protocols also differed in solution contact time, volume of leachate, presence or absence of agitation, additional filtration step using a backing filter, and pore size of any backing filter, as summarized in Table 2. Thus, it is important to note that, even within each of the three categories of methods, no leaching protocol examined in this study was identical to any other.

Results are presented hereafter using the acronym of the laboratory followed by the type of leaching solution used ("-u" for UPW, "-a" for AmmAc and "-b" for Berger). Where groups performed sequential leaching (Table 2), we consider the sum of both leaching fractions to be equivalent to a single-step AmmAc (for UTAS-a) or Berger (for UGA-b) leach.

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

The digestion procedure used by the GIG laboratory to measure total trace elements contained in aerosols was described previously (Zhang et al., 2022). Briefly, aerosol filter subsamples were digested in a mixture of HNO_3, H_2O_2 and HF in an acid-cleaned Teflon jar, using microwave digestion. The residual solution was then evaporated, and 20 mL HNO_3 (1%) was added to the jar. The resulting solution was filtered through a 0.22 μm polyethersulfone filter and analyzed using inductively coupled plasma mass spectrometry (ICP-MS, iCAP Q, Thermo Fisher Scientific).

205 **Table 2.** Overview of the eight leaching protocols used by the six laboratories participating in this intercomparison study.

protocol	batch/flow-through	contact time	leaching solution	volume	agitation	filter pore size (µm)	reference
GIG	batch	120 min	ultrapure water (pH: 6.5)	20 mL	orbital shaking	0.22	Zhang et al., 2022
NIO	batch	30 min	ultrapure water (pH: 6.4)	20 mL	ultrasonication	0.2	Panda et al., 2022
UEA	batch	60 min	ammonium acetate (1.1 mol/L, pH: 4.7)	20 mL	hand shaking	0.2	Sarthou et al., 2003
UoP	flow-through	~45 s	ultrapure water (pH: 5.2)	100 mL	no agitation	0.2	Buck et al., 2010
UGA-leach 1	flow-through	~20 s	ultrapure water (pH: 5.6)	100 mL	no agitation	0.2	Buck et al., 2013
UGA-leach 2 (following leach 1)	batch	~120 min	25% acetic acid (v/v) + 0.02 mol/L hydroxylamine hydrochloride	10 mL	no agitation	no backing filter (centrifugation)	Berger et al., 2008
UTAS-leach 1	flow-through	~40 s	ultrapure water (pH: 6.5)	50 mL	no agitation	no backing filter	Perron et al., 2020a
UTAS-leach 2 (following leach 1)	batch	60 min	ammonium acetate (1.1 mol/L, pH: 4.7)	10 mL	hand shaking	no backing filter (centrifugation)	Perron et al., 2020a

206

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

2.4 Trace element determination

Each group determined trace elements in the leaching solutions using the analytical methods routinely used in their laboratory. Table 3 shows the trace elements determined by each laboratory, Table S7 summarizes the methods used, and Table S8 provides a summary of the analytical detection limits. As shown in Table 3, up to 20 trace elements were determined by groups participating in this intercomparison study. The results presented in this paper focus on the seven trace elements (i.e., Al, Cu, Fe, Mn, Ni, Pb and V) which were determined by all the six groups. Comparisons for the other elements (Ba, Cd, Ce, Co, Cr, La, P, Th, Ti, U and Zn) are reported in the supplementary material (Figures S1 and S3-S6). For all groups, the mass of trace elements in aerosol samples was determined against an external standard calibration (Table S7). The same trace elements determined in leaching solutions by GIG were also determined in the total digests of the Guangzhou and Qingdao samples. Trace elements concentrations in leachate and digests were corrected for the volume of the liquid phase used and reported as mass of trace elements on the filter subsamples received.

Table 3. List of trace elements determined by each of the six laboratories. If one trace element was not measured by one laboratory, the corresponding cell is left blank.

Trace element	GIG	NIO	UEA	UTAS	UGA	UoP
Al	y	y	y	y	y	y
As	y		y			
Ba	y		y	y		y
Cd	y		y	y	y	y
Ce			y	y		
Co		y	y	y	y	y
Cr	y		y	y	y	y
Cu	y	y	y	y	y	y
Fe	y	y	y	y	y	y
La			y	y		
Mn	y	y	y	y	y	y
Ni	y	y	y	y	y	y

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

Pb	y	y	y	y	y	y
P		y	y	y		
Sb	y		y			
Th		y	y	y		
Ti		y	y	y	y	y
U			y	y		y
V	y	y	y	y	y	y
Zn	y		y	y	y	y

2.5 Blank subtraction

For each analytical method, blank subtraction was conducted using the following procedure.

As described above, each group received 13 blank filters (B1-B13). First, we calculated the median of all the blank measurements generated by a research group. If the median value was below the analytical detection limit for that trace element and analytical method, no blank correction was made; if the median was above the analytical detection limit, this median value was subtracted from the respective measured quantity for the subsample. We calculated the median absolute deviation (MAD, defined as the median of the absolute differences between the individual blanks and the median blank) to represent the uncertainty in the blank. Subsamples with blank-corrected quantities less than three times of the MAD were defined as less than the blank-correction detection limit. In this work we use median and MAD, rather than mean and standard deviation, because median and MAD can be reliably calculated when values below detection limits are present and they are less affected by the presence of outliers.

2.6 Data analysis

2.6.1 Sample homogeneity

Six samples were collected in Guangzhou to examine particle distribution homogeneity, and eight subsamples (47 mm discs) were obtained from each sample. For each filter sample, we measured the mass of 14 trace elements in the eight subsamples (X_i, where X_i is the mass

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

of a given trace element in the *i*th subsample) at the GIG laboratory; we then calculated the median value (X_m) of the eight measurements, the MAD as well as relative MAD, which is defined as MAD/X_m.

2.6.2 Precision derived from replicate subsample analysis

Each group measured trace elements on three subsamples from each of two different aerosol samples (C samples) collected at Qingdao (Table 1). For each trace element, the relative MAD was determined for both samples and the higher of these values was used as the uncertainty in the soluble trace element mass measurement. This value was then applied to the intercomparison study samples (D1-D26) for which no replicate sample was available. We note that the uncertainty determined in this way includes a component due to the heterogeneity of trace element distribution across the aerosol samples, as well as a component from variability in the analytical procedures within each group.

For four (GIG, NIO, UGA and UTAS) of the six groups participating this work, the magnitude of the uncertainty in each intercomparison subsample was determined by multiplying its measured soluble mass by its respective relative MAD value. The other two groups (UEA and UoP) reported individual uncertainties for the intercomparison subsamples based on calibration uncertainties and repeated analyses of single subsample extract solutions. In the latter cases, the uncertainties of the soluble trace element mass measurement in individual subsamples were taken to be whichever of the replicate-determined or individually-determined uncertainties was larger.

2.6.3 Inter-method statistical comparisons

Comparison of the results produced by the various leaching methods for each trace element for samples D1 – D26 was done where possible using both parametric and non-parametric tests, and each dataset was tested for normality using the Shapiro Wilks method (Miller and Miller, 2010). For comparisons of similar methods (e.g., the batch UPW leaches of

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

GIG and NIO) the hypothesis that each method-method pair produced statistically indistinguishable results was examined according to the following statistical tests.

(1) Pearson's and Spearman's Rank Correlations were used to test the correlation between methods (assuming that there is a direct relationship between the amounts of a trace element leached by the two methods). When no correlation was found between two methods, subsequent tests were deemed to be unreliable.

(2) A two-tailed *t*-test was used to test the slope of a method-method relationship. A slope equal to 1 implies that sample-to-sample changes in trace element release are equivalent between the two methods. Because the samples shared for the intercomparison were subject to heterogeneity, the slope was also tested for difference to 1±0.12, after uncertainties due to sample heterogeneity were taken into account (the upper limit of such uncertainties is estimated to be 12%, as discussed in Sections 2.6.1 and 3.1); in this case, one-tailed *t*-tests were used.

(3) A two-tailed *t*-test was used to investigate whether the intercept of the method-method relationship differed significantly from zero. Divergence from zero indicates the presence of an offset between the two methods.

(4) Paired *t*-tests and Wilcoxon Signed Rank tests were used to assess if the difference between the two measurements for each individual sample was different from zero.

Optimum agreement between each method-method pair was therefore indicated by significant correlation (test 1) and absence of significant differences indicated by tests 2-4.

Comparisons between dissimilar methods (e.g. batch UPW vs AmmAc) were made using two-tailed t-tests and Mann Whitney U tests, after first confirming the presence of significant differences between the datasets using one-way ANOVA and Kruskal Wallis tests.

2.7 Air mass back trajectory analysis

Five-day air mass back trajectories (AMBTs) were calculated at heights of 50, 500 and 1000 m above the ground level throughout the sampling period at the sampling site in Qingdao,

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

using the NOAA READY HYSPLIT model with NCEP/NCAR Reanalysis Project datasets (Stein et al., 2015). For each sample, trajectories were calculated every 3 hours throughout the collection period and 50 m trajectories (for all samples) were used to perform cluster analysis with the openair software tool in R (Carslaw and Ropkins, 2012). Between four and six clusters were tested for this analysis, and three clusters (arrivals from the north (N), southwest (SW) and the Yellow Sea (YS)) were chosen to represent the major differences in atmospheric transport pathways during the field sampling (Figure 1). Each sample was assigned to one of the three clusters, based on the dominant AMBT type of the eight trajectories calculated for that sample.

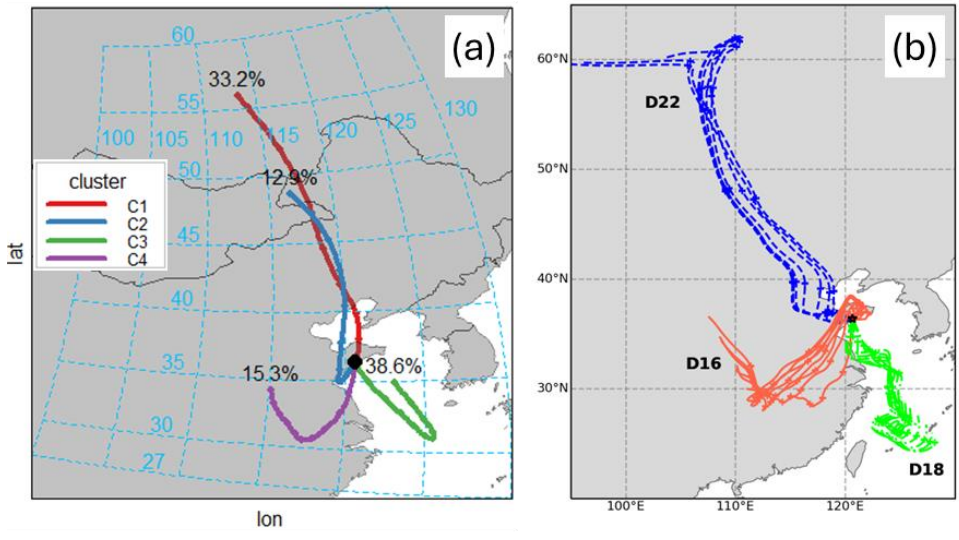


Figure 1. (a) Result of the 4-cluster analysis of the AMBTs calculated during the sample collection at Qingdao, showing representative pathways for the clusters and their percentage occurrence; (b) examples of 50 m AMBTs for samples classified as N (Sample D22, blue dash), SW (D16, red solid) and YS (D18, green dot-dash) types.

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

3 Results and Discussion

3.1 Aerosol trace element distribution on the filters

Table 4 summarizes the relative MAD values obtained for total trace elements contained in the Guangzhou samples. The median of relative MAD values obtained for the 14 trace elements ranged from a minimum of 2.7% for V to a maximum of 12% for Cr. In this work, in order to disentangle subsample variability from analytical method variability we chose a conservative approach by applying the highest median relative MAD value (i.e. 12%) to represent the uncertainty associated with trace element distribution heterogeneity over a filter, regardless of the element analyzed.

Table 4. Summary of the relative MAD (in %) in particle distribution homogeneity: minimum, maximum and median values for the 6 filters collected in Guangzhou.

Element	Min	Max	Median
Al	1.4	11	7.1
As	1.0	8.0	5.1
Ba	1.4	12	5.9
Cd	0.5	10	5.0
Cr	6.6	22	12
Cu	2.1	14	4.2
Fe	1.6	11	4.9
Mn	1.1	12	6.5
Ni	3.4	11	5.4
Pb	0.6	7.7	4.5
Se	3.5	10	6.9
Sb	1.4	8.9	4.8
V	1.3	9.4	2.7
Zn	3.2	10	8.7

3.2 Background (blank) contribution

Each research group determined the soluble fraction of trace elements in blank samples

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

(B samples) according to their chosen method(s); therefore, the blank values included the contribution from the blank filters and from the analytical procedures used. The total blank for each trace element analyzed was subtracted from the subsamples' measurements in the intercomparison study, without attempting to separately quantify the contributions from the filter and the analytical procedure.

In almost all cases, blank contributions were extremely low with respect to the PM_{10} subsamples. For the seven elements that are the focus of this intercomparison study, 79% of blanks were below the analytical detection limit, and only two blank values were greater than 10% of the magnitude of the lowest intercomparison sample. There were only a few cases for which blanks were higher than the lowest intercomparison sample and 24 cases (~2%) where intercomparison samples were identified as being below the blank-correction detection limit. Table S9 summarizes the blank contributions for all trace elements with measurable blanks.

3.3 Intra-method precision

Table 5 summarizes the relative uncertainties (in %) for soluble trace elements determined by each of the six groups when analyzing three replicates of the same aerosol sample (C samples), and the relative uncertainty is defined in Section 2.6.2. In most cases these relative uncertainties were within the range of values determined for total element homogeneity over whole filter samples (Table 4). None of the uncertainties reported in Table 5 was larger than the highest uncertainty value reported for the total element homogeneity test (22% for Cr), suggesting that intra-laboratory variability was not significantly greater than the variability between subsamples within individual filter samples.

346 **Table 5.** Summary of the relative uncertainties (in %) for soluble trace elements analysis of 3 replicate samples determined by each of the six
347 laboratories. If one trace element was not measured by one laboratory, the corresponding cell is left blank.

Trace element	GIG-u	NIO-u	UEA-a	UTAS-u	UTAS-a	UGA-u	UGA-b	UoP-u
Al	11	2.6	4.4	18	15	9.1	11	8.2
As	11		3.2					
Ba	4.5		2.9	18	4.2			4.9
Cd	5.7		2.8	10	8.4	0.11	1.5	7.0
Ce			0.8	14	11			
Co		7.8	4.3	4.7	0.9	1.0	3.6	4.4
Cr	12		1.1	2.7	4.1	1.2	1.8	21
Cu	6.9	6.8	15	3.8	6.1	16	16	11
Fe	12	9.0	1.6	17	15	2.0	2.0	13
La			2.3	17	8.3			
Mn	9.7	1.7	4.7	6.1	5.1	3.6	4.9	3.7
Ni	15	6.9	5.0	2.6	7.4	8.5	1.5	15
P		5.3	1.1	0.6	4.7			
Pb	12	6.9	3.5	8.2	4.7	2.0	3.4	10
Sb	7.8		2.5					
Th		12	0.6	5.8	15			
Ti		7.2	7.9	25	7.3	0.12	43	5.6
U			2.1	11	9.5			8.1
V	10	4.8	3.1	5.1	1.8	11	3.0	9.6
Zn	12		4.7	15	6.3	7.3	7.3	5.2

348

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

3.4 Comparison of leaching methods

Figure 2 summarizes statistical comparisons of similar leaching methods for all of the samples collected at Qingdao (D samples), and the results are discussed below.

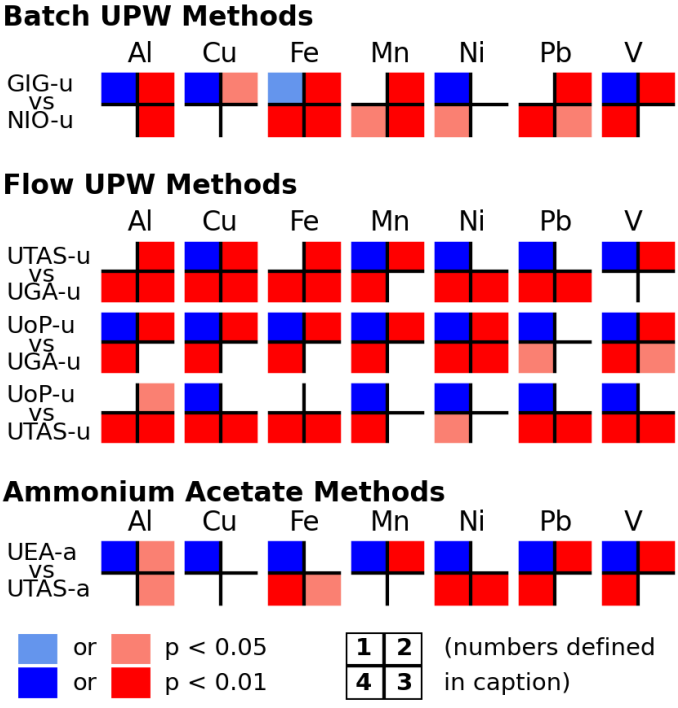


Figure 2. Summary of the statistical comparisons of the leaching methods for all of the 26 samples (N, SW and YS) collected at Qingdao. Comparisons are only shown for similar methods: Batch UPW (GIG-u and NIO-u), Flow UPW (UoP-u, UTAG-u and UGA-u) and AmmAc (UEA-a and UTAS-a). Statistical tests (as detailed in Section 2.6.3) shown in quadrants are (1) Spearman's Rank Correlation, (2) one-tailed t-test of slope versus 1±0.12 uncertainty, (3) intercept equals zero, and (4) Wilcoxon Signed Rank Test. Significant test results are indicated by the color code: blue indicates agreement between methods, red indicates disagreement, and white indicates that the test result was not statistically significant. Full statistical test results (all tests and elements) can be found in Figure S4.

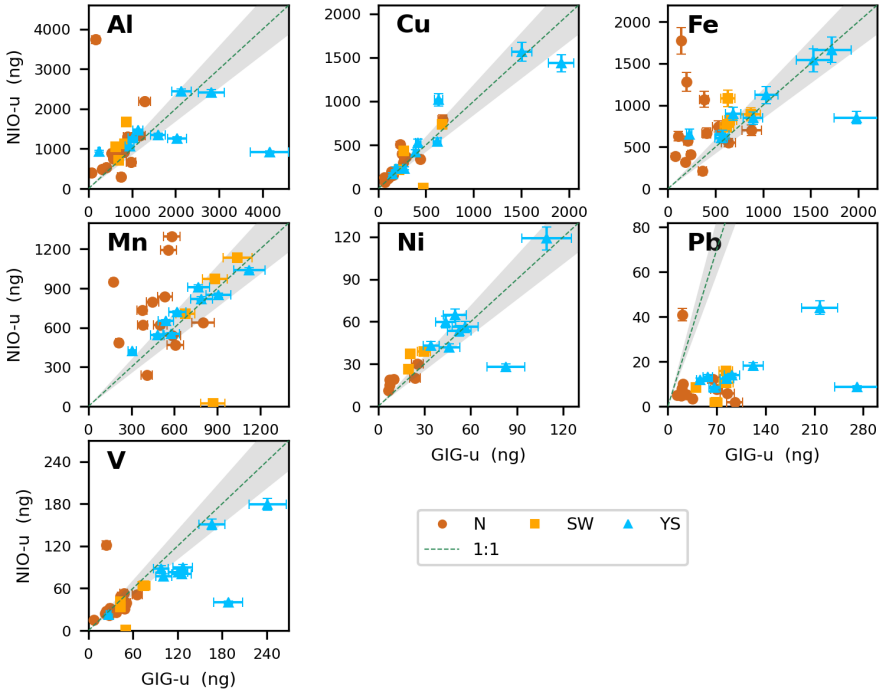
https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

362

363 **3.4.1 Ultrapure water batch leaching methods**

364 Table 2 shows that the two UPW batch leaching protocols examined in this work differ in
365 contact time (2 versus 0.5 h), agitation method (orbital shaking versus ultrasonication), and to
366 a lesser extent, filter pore size (0.22 versus 0.2 μm).

367 As shown in Figure 2, a strong correlation (Spearman test, $p<0.01$) was observed between
368 soluble trace element masses measured by GIG and NIO for four of the elements investigated
369 (Al, Cu, Ni and V); nevertheless, the slope statistically differs from 1±0.12 for Al, Cu and V
370 and the intercept is statistically different from zero in the case of Al. Among these four elements,
371 the best agreement between the two protocols was found for Cu and Ni (Figure 3), with NIO
372 values slightly higher than those of GIG, except for a few YS samples. Good agreement was
373 also observed for Al and V, although for both elements there are some samples with measured
374 values greatly deviating from the 1:1 line.



375

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

Figure 3. Comparison of the absolute mass of soluble trace elements (Al, Cu, Fe, Mn, Ni, Pb and V) obtained using UPW batch leaching methods (GIG and NIO). Air mass origins (N, SW, YS) for each sample are indicated by the color code. The dashed line indicates the 1:1 relationship between the methods, and the grey shading indicates the 12% sample homogeneity uncertainty.

Although there was a significant correlation for Fe (Spearman test, $p<0.05$, Figure 2), large deviation from the 1:1 slope was not uncommon (Figure 3). Fe measured by NIO tended to be higher than GIG values, especially for samples from continental air masses (N and SW). No significant correlation was found between the two methods for Mn or Pb (Figure 2). Similar to Fe, soluble Mn values reported by NIO tended to be higher than those reported by GIG (Figure 3), especially for N and SW samples. The ultrasonication used by NIO may lead to the formation of reactive oxygen species (e.g., hydroxyl radicals and hydrogen peroxide) due to acoustic cavitation (Kanthale et al., 2008; Miljevic et al., 2014). These species have the potential to reduce insoluble Fe(III) and Mn(IV) to the more soluble Fe(II) and Mn(II), especially when poorly soluble mineral dust is present in the N and SW samples. Enhanced solubility of Al in the NIO data (relative to GIG) is not observed, perhaps because Al only has one oxidation state and its solubility may not be affected by redox chemistry.

Notably, NIO reported much lower values than GIG for Pb (Figure 3). GIG and NIO both checked their experimental data and found no errors, and there is no clear clue why such large differences occurred. Therefore, here we do not discuss the UPW batch data for Pb further.

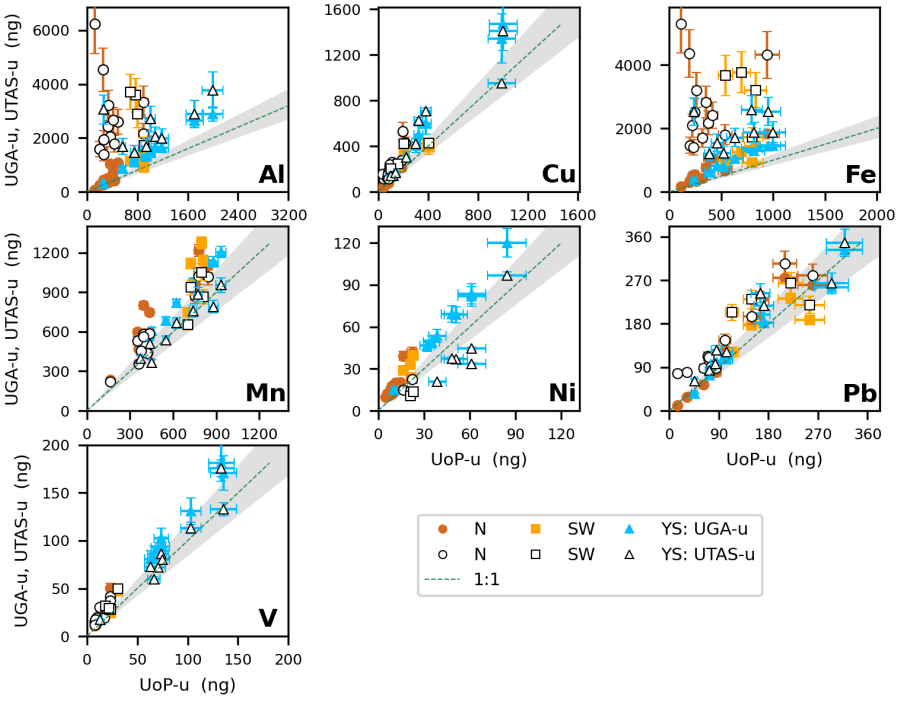
3.4.2 Ultrapure water flow-through leaching methods

The three UPW flow-through leaching methods investigated here differ by contact time, pH and volume of leaching solution, and the use (or not) of a backing filter (Table 2).

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

Overall, a positive and significant correlation (Spearman test $p<0.01$, Figure 2) was found between the soluble trace element mass measured by UoP and UGA. For all trace elements except Pb, the comparison of these two methods, however, resulted in a slope different from 1±0.12 and an intercept different from 0 (Figure 2), with slightly lower values obtained by UoP than by UGA (Figure 4). This difference cannot be attributed to higher blank levels in UGA samples because both laboratories had element blanks below their method detection limits (hence, no blank correction was applied).

The UTAS UPW flow-through method produced soluble trace element measurements that positively and significantly correlate with the two other methods (Figure 2), except for Al and Fe. Overall, UTAS measurements were higher than those of UoP (Figure 4), except for low Ni measurements (compared to both UoP and UGA) which result from elevated Ni blank correction in UTAS data treatment (due to new Ni cones used in the SF-ICP-MS instrument, Table S9).



https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

Figure 4. Comparison of the absolute mass of soluble trace elements (Al, Cu, Fe, Mn, Ni, Pb and V) obtained using UPW flow-through leaching methods (UoP, UTAS and UGA). Plot details are as described in Figure 3. UGA data are represented by solid symbols, and UTAS data are represented by open symbols.

One striking observation from Figures 2 and 4 is the lack of correlation between Al and Fe measured by UoP or UGA and that measured by UTAS (similar behavior was also observed for Ti, Ba and U, as shown in Figure S1). High Al and Fe masses reported by UTAS could stem from the absence of a backing filter in the UTAS method, whereby the soluble fraction of trace elements measured can include the contribution of particles with a diameter greater than 0.2 µm. As shown in Figures 4 and S2, this observation is even more obvious in aerosol samples influenced by terrestrial air-masses (N and SW samples) due to lower solubility of Fe and Al in these samples when compared to YS samples. The absence of a backing filter in the UTAS protocol only seems to influence lithogenic elements with lower solubility (Al, Fe, Ti, Ba, U), while more soluble elements (Cu, Mn, Ni, Pb and V) showed a significant ($p<0.01$) correlation with the other two methods investigated. No filtration in the UTAS UPW flow-through protocol may also be responsible for higher uncertainty obtained for lithogenic elements when measuring three replicate samples (Table 5).

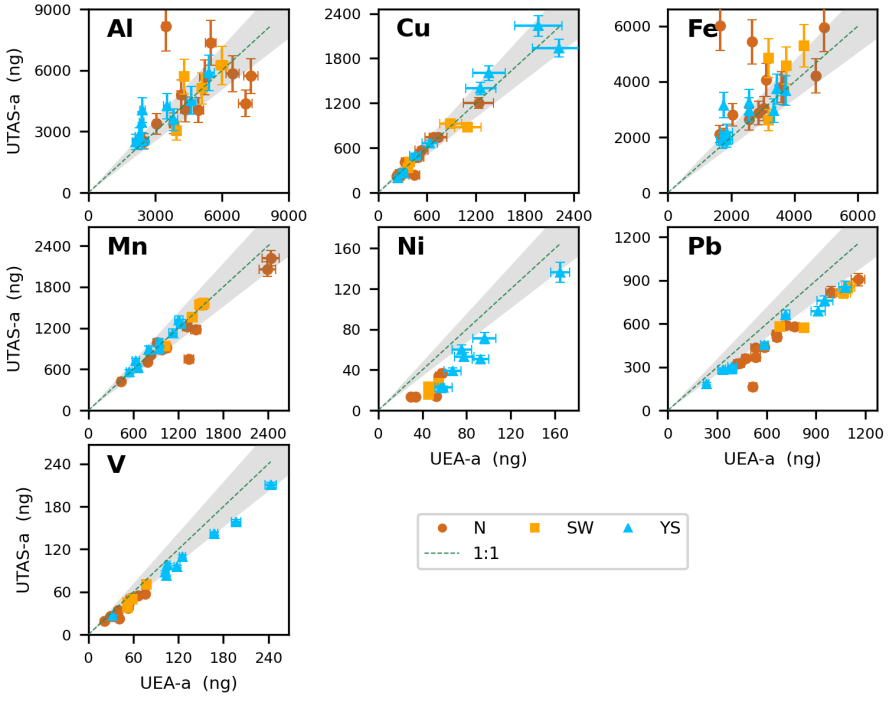
Based on the intercomparison results, it is reasonable to assume that the different leaching solution volumes have minor, if any, impacts on the results from the three UPW flow-through leaching methods. Our findings are consistent with those of Winton et al. (2015) suggesting that over 90% of the soluble Fe contained in aerosols was extracted using a single 50 mL UPW flow-through leach, although Winton et al. (2015) did not provide for other trace elements.

3.4.3 Ammonium acetate leaching methods

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

As displayed in Table 2, the two AmmAc leaching protocols applied by UEA and UTAS differ by the volume of the leaching solution and the absence of a backing filter in the UTAS protocol; in addition, the UTAS leaching is performed as part of the sequential leaching of a single sample, immediately following the UPW flow-through leaching.

For Cu and Mn (to a lesser extent), measurements show very good agreement (significant correlation ($p<0.01$) and few significant differences in slopes, intercepts and soluble masses) for the two AmmAc methods (Figures 2 and 5). Similarly to the UPW flow-through comparison (Section 3.4.2), there was an offset between the two Ni datasets, possibly due to the Ni blank overcorrection applied to the UTAS dataset (see Section 3.4.2 and Table S9) which shifts the correlation curve below the 1±0.12 grey-shaded area in Figure 5. Both Pb and V measurements show good correlation and agreement in intercept values, although the slopes differ from 1±0.12 (Figure 2). Diverging calibration between the two methods may explain such discrepancies.



https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

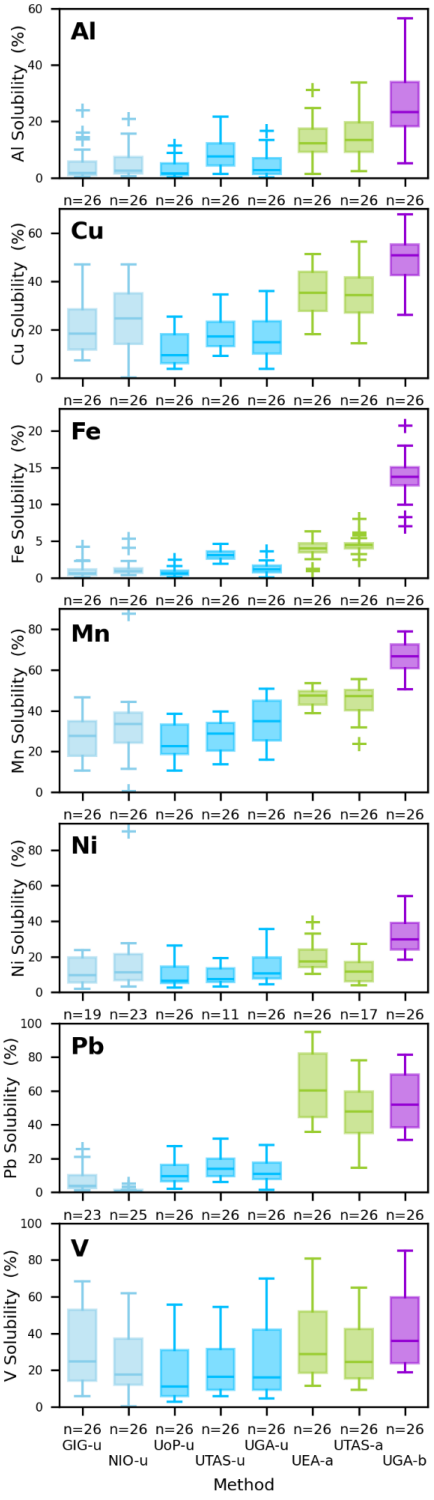
Figure 5. Comparison of the absolute mass of soluble trace elements (Al, Cu, Fe, Mn, Ni, Pb and V) obtained using AmmAc extraction methods (UEA and UTAS). Plot details are described in Figure 3.

Poorer agreement between the two methods was again found for lithogenic elements (Al and Fe), although comparisons were noticeably better for the AmmAc methods (Figure 5) than for the UTAS-u to other UPW flow-through method comparisons (Figure 4). For these two elements (Al and Fe), measurement differences were more pronounced in samples influenced by "terrestrial" air masses (N and SW) while YS samples showed a good agreement (Figures 5, S5 and S6). This better agreement for lithogenic elements between the AmmAc methods may suggest that the > 0.2 μm particles which are not removed by the UTAS UPW leaching protocol are largely soluble in AmmAc so that their trace element content is not removed by 0.2 μm filtration in the UEA method.

3.4.4 Influence of different leaching protocols on the determination of fractional solubility

Figure 6 illustrates the variations in solubility of trace elements in the Qingdao samples, determined using the eight leaching methods tested in this study. Kruskal Wallis and one-way ANOVA tests both indicated significant differences ($p < 0.01$) among the results obtained for the five methods employing UPW as the leaching solution for all elements except Ni and V. Some of the significant differences within the UPW methods are discussed in Sections 3.4.1-3.4.2 (e.g. for Pb in the UPW batch methods, and Al and Fe in the UPW flow-through methods). In general, where significant differences in solubility existed within the UPW group these did not appear to be related to whether the method employed batch or flow-through techniques.

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025



474

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

Figure 6. Box and whisker plots of solubility (in %) for Al, Cu, Fe, Mn, Ni, Pb and V in the Qingdao samples determined using the eight leaching protocols tested in this study. Colors indicate method type: batch UPW (light blue), flow UPW (dark blue), AmmAc (green), and acetic acid with hydroxylamine hydrochloride (purple).

When comparing all the eight methods, Kruskal Wallis and one-way ANOVA tests indicated significant differences in solubility for all seven elements. In almost all cases, the UPW methods resulted in significantly different (lower) solubilities than those obtained with the AmmAc and Berger leaches, as expected. For Al, Cu, Fe and Mn, solubilities obtained using the AmmAc leach were significantly different (lower) than those obtained using the Berger leach.

Table 6 illustrates the broad differences obtained between the UPW, AmmAc and Berger methods using the pooled median solubility values for each method. Both Figure 6 and Table 6 show that the extent of solubilization from aerosol by the different methods is highly element specific. For example, the patterns in enhancement in solubility for AmmAc and Berger compared to UPW (AmmAc/UPW and Berger/UPW, see Table 6) are very different for Fe, Mn and Pb. Interestingly, the more aggressive conditions of the Berger leach (compared to AmmAc) had large impacts on the dissolution of poorly soluble lithogenic elements (such as Al and Fe), but no effect on Pb dissolution.

Table 6. Median trace element solubilities for the Qingdao aerosol samples, determined from the combined soluble element masses determined for the methods using UPW (GIG-u, NIO-u, UoP-u, UTAS-u and UGA-u), AmmAc (UEA-a and UTAS-a) and Berger (UGA-b), and solubility ratios of the AmmAc to UPW (AmmAc/UPW) and Berger to UPW (Berger/UPW) methods. (a: excluding UTAS-u due to the inclusion of > 0.2 μm particles, b: excluding GIG-

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

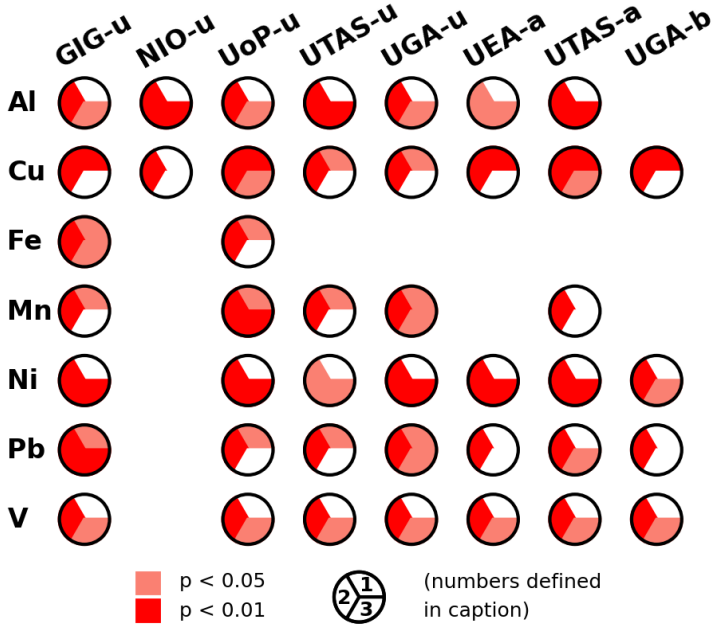
u and NIO-u due to unexplained inconsistencies in Pb data).

element	UPW Solubility (%)	AmmAc Solubility (%)	Berger Solubility (%)	AmmAc / UPW	Berger / UPW
			All samples		
Al	2.1 [a]	13	23	6.1	11
Cu	16	34	51	2.2	3.2
Fe	0.9 [a]	4.2	14	4.7	15
Mn	29	47	67	1.6	2.3
Ni	8.2	14	30	1.7	3.6
Pb	8.9 (11 [b])	57	52	6.4 (5.0 [b])	5.9 (4.6 [b])
V	16	25	36	1.5	2.2
			N samples		
Al	1.3 [a]	12	20	9.1	15
Cu	11	27	42	2.4	3.7
Fe	0.6 [a]	4.0	14	7.0	24
Mn	20	42	61	2.1	3.0
Ni	6.0	10	29	1.7	4.7
Pb	6.0 (8.4 [b])	42	39	7.1 (5.1 [b])	6.4 (4.6 [b])
V	11	19	28	1.7	2.5
			SW samples		
Al	1.9 [a]	11	21	5.5	11
Cu	19	39	55	2.1	2.9
Fe	0.9 [a]	4.1	14	4.8	16
Mn	30	46	66	1.5	2.2
Ni	6.6	12	24	1.8	3.7
Pb	10 (11 [b])	54	54	5.6 (5.0 [b])	5.7 (5.0 [b])
V	16	23	31	1.4	2.0
			YS samples		
Al	8.2 [a]	19	37	2.3	4.5
Cu	28	45	57	1.6	2.0
Fe	1.6 [a]	4.4	15	2.8	9.6
Mn	38	49	71	1.3	1.9
Ni	19	22	45	1.2	2.4
Pb	15 (19 [b])	73	76	4.9 (3.8 [b])	5.1 (3.9 [b])
V	42	52	62	1.2	1.5

For most methods, there were statistically significant differences (Kruskal Wallis and one-way ANOVA tests) between solubilities determined for the different air mass types for most

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

505 elements (Figures 7 and S7). Exceptions to this were the NIO-u results (for which significant
506 differences were only found for Al and Cu) and Fe (for which significant differences were
507 found for some of the UPW methods, but not for AmmAc or Berger). In all cases where such
508 significant differences were found, solubility in YS-type samples was higher than that in N-
509 type samples (Figure 7). Note that the difference between N and YS samples appear to become
510 less pronounced as the leach solution becomes more aggressive (e.g. for V, the YS/N solubility
511 ratios are 4.0-6.1 for the UPW methods, 2.7-3.0 for the AmmAc methods, and 2.3 for the
512 Berger method). Enhanced trace element solubility in YS samples compared to N samples may
513 be due to increased atmospheric processing of particles transported over the ocean (particularly
514 for lithogenic elements) or anthropogenic emissions of highly soluble trace elements to the
515 marine atmosphere (e.g. V, Ni from shipping).



516

517 **Figure 7.** Summary of the comparisons of trace element solubility for the N, SW and YS
518 samples determined using the eight leaching methods tested in this work. Red colors indicate
519 statistically significant differences (Mann Whitney U Test) between pairs of sample types (1:

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

N vs SW; 2: N vs YS; 3: SW vs YS), and white colors indicate no significant difference. No summary is shown where the Kruskal Wallis Test indicated no significant differences ($p < 0.01$) between the three sample types.

The aerosol provenance seems to be a key driver of the resulting amount of trace element measured using the UPW flow-through leach. Indeed, higher Al, Cu, Ni, and V solubility was determined for YS samples compared to samples showing terrestrial fingerprints. Such an increase in trace element solubility in marine samples was less pronounced when measuring Fe, Mn and Pb due to similarly high solubility measured in samples influenced by SW air-masses. In YS samples, V, Ni and Cu showed higher solubility in UPW, with a lesser impact of stronger AmmAc and Berger leaching protocols. Increased atmospheric processing of particles transported over the ocean can explain the presence of more readily soluble trace elements in the presence of UPW in YS samples compared to N and SW samples.

As opposed to the UPW leaches, AmmAc methods showed similar (Mn, Pb) or higher (Al, Fe) solubilities in samples containing terrestrial inputs, except for Cu, Ni and V for which higher solubility was found in samples influenced by marine air masses.

4 Conclusions and recommendations for future studies

In general, comparisons between similar leaching methods (Sections 3.4.1-3.4.3) show a high degree of correlation between measured soluble masses of trace elements. Where correlation was poor, the differences appear to stem from specific differences in the leaching methods. For example, the use of ultrasonic agitation instead of mechanical agitation for UPW batch methods seems to result in increased solubility for Fe and Mn, and the absence of a backing-filter resulted in higher soluble Al and Fe in the UPW flow-through leach. The trace elements most impacted by these experimental differences appear to be lithogenic elements associated with mineral dust, including Fe.

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

Most of the elements whose soluble masses are well-correlated nevertheless show significantly different slopes and intercepts from those which would be expected for "identical" results (within the uncertainties associated with sample heterogeneity). Differences in calibration between the participating research groups, as well as the differences in leaching procedures (Table 2), could contribute to this behavior. All groups took steps to determine the accuracy of their calibrations using external reference materials (Table S7), but there was no common procedure adopted during this intercomparison exercise to verify the accuracy of analysis, suggesting a common approach and sharing of certified reference materials (CRMs) could further improve confidence in the results of analytical intercomparisons.

The differences in fractional solubility observed between the different leaching method types in this study are not unexpected, since dissolution of trace elements is partially dependent on the chemical characteristics of the leach solution (which varies widely between UPW, AmmAc and Berger). The GEOTRACES data products naming convention (https://www.geotraces.org/parameter-naming-conventions/; last accessed: 18 June 2025) suggests describing both AmmAc and Berger leaches as "strong" leaches. However, for some elements studied here, there is a wide range in trace element solubility data reported within this suggested classification. For Pb and V, the AmmAc and Berger leaches appear to give equivalent results, while solubility determined using AmmAc and Berger leaches was significantly different for Al, Cu, Fe, Mn and Ni. Careful consideration should be given to the variation in solubility between leach solutions for a given trace element when observational data are used to validate numerical models, since using data obtained from a variety of analytical methods might introduce unknown biases into the validation.

The differences in trace element solubility observed for aerosols with different source and/or transport histories within each of the individual leaching methods further complicates this issue and highlights the importance of underlying particle properties in determining trace

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

element solubility. While Fe (the most commonly studied and modelled trace element) appears to be least affected by differences in sample source or transport type for the trace elements reported here, it would appear to be necessary to investigate such behavior in aerosols from a wider variety of environments and regions in order to fully quantify this effect.

Our study and others (e.g. Perron et al., 2020a; Li et al., 2023) clearly demonstrate that different leaching solutions release different proportions of aerosol trace elements into solution. It is not possible to identify a "correct" procedure for the determination of aerosol soluble element input to the ocean (or indeed whether a correct procedure exists) without a better understanding of the factors that control trace element dissolution in the complex and variable matrix of natural seawater over the time frames during which aerosol particles remain suspended following deposition to the ocean (Baker and Croot, 2010). The identification in future work of links between different leaching protocols and specific environmental processes or behaviors would be a significant advance.

However, this study highlights the necessity for some best practice guidance to reduce uncertainties in future intercomparison studies of aerosol soluble trace element leaching protocols. Best practice should include an agreed approach to analytical instrument calibration, representative blank definition and detection limit determination. Distribution of one or more solution-phase reference and/or consensus samples relevant to the range of trace elements assessed would also be advisable. Reporting by each group of measured concentrations for these solution-phase samples would allow direct intercomparison of analytical performance that is independent of extraction methods. In addition, the reporting of measurement uncertainty and precision via a common approach should be encouraged for all future studies. Adoption of such best practices outside of intercomparison studies will improve intercomparability of aerosol soluble trace element data, helping to improve the robustness of data interpretation and assimilation of observational data into models.

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

Data availability.

Data used in this paper are available from https://doi.org/10.5281/zenodo.15728006.

Competing interests.

At least one of the (co-)authors is a member of the editorial board of Atmospheric Measurement Techniques.

Author contribution.

Conceptualization: Mingjin Tang, Morgane M.G. Perron, Alex R. Baker, Rui Li, Andrew R. Bowie, Clifton S. Buck, Ashwini Kumar, Rachel Shelley, Simon Ussher; **Aerosol sampling:** Mingjin Tang, Rui Li; **Laboratory analysis:** Rui Li, Rachel Shelley, Rob Clough, Scott Meyerink, Prema P. Panda, Ashley T. Townsend, Neil Wyatt; **Data analysis:** All authors; **Statistical analysis:** Alex R. Baker, Rui Li; **Writing - Original Draft:** Mingjin Tang, Morgane M.G. Perron, Alex R. Baker, Rui Li, Andrew R. Bowie, Clifton S. Buck, Ashwini Kumar, Rachel Shelley, Simon Ussher; **Writing - Review & Editing:** all the authors; **Funding acquisition:** Mingjin Tang, Alex R. Baker, Andrew R. Bowie, Clifton S. Buck, Ashwini Kumar, Simon Ussher.

Financial support.

Mingjin Tang was funded by National Natural Science Foundation of China (42321003 and 42277088), International Partnership Program of Chinese Academy of Sciences (164GJHZ2024011FN), Guangzhou Bureau of Science and Technology (2024A04J6533), and Guangdong Foundation for Program of Science and Technology Research (2023B1212060049). Alex R. Baker and Rachel Shelley were funded by the UK Natural Environment Research Council (grant number NE/V001213/1). Morgane M.G. Perron was funded by a European Marie Sklodowska-Curie Actions fellowship (GA 101064063). Andrew R. Bowie was funded by the Australian Research Council (DP190103504) and the Australian

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

Antarctic Program Partnership under the Antarctic Science Collaboration Initiative (ASCI000002). This work is also partly funded by Scientific Committee on Oceanic Research (SCOR) Working Group 167 (Reducing Uncertainty in Soluble aerosol Trace Element Deposition, RUSTED).

Acknowledgement.

We would like to thank members of the Scientific Committee on Oceanic Research (SCOR) Working Group 167 (Reducing Uncertainty in Soluble aerosol Trace Element Deposition, RUSTED) for valuable discussion and colleagues at Shandong University for their assistance in aerosol sampling.

References

Aguilar-Islas, A., Planquette, H., Lohan, M. C., Geibert, W., and Cutter, G.: Intercalibration: A Cornerstone of the Success of the GEOTRACES Program, Oceanography, 37, 21-24, 2024.

Bai, X., Tian, H., Zhu, C., Luo, L., Hao, Y., Liu, S., Guo, Z., Lv, Y., Chen, D., Chu, B., Wang, S., and Hao, J.: Present Knowledge and Future Perspectives of Atmospheric Emission Inventories of Toxic Trace Elements: A Critical Review, Environ. Sci. Technol., 57, 1551-1567, 2023.

Baker, A. R. and Croot, P. L.: Atmospheric and marine controls on aerosol iron solubility in seawater, Mar. Chem., 120, 4-13, 2010.

Baker, A. R. and Jickells, T. D.: Mineral particle size as a control on aerosol iron solubility, Geophys. Res. Lett., 33, L17608, doi: 17610.11029/12006GL026557, 2006.

Baker, A. R., Kanakidou, M., Nenes, A., Myriokefalitakis, S., Croot, P. L., Duce, R. A., Gao, Y., Guieu, C., Ito, A., Jickells, T. D., Mahowald, N. M., Middag, R., Perron, M. M. G., Sarin, M. M., Shelley, R., and Turner, D. R.: Changing atmospheric acidity as a modulator of nutrient deposition and ocean biogeochemistry, Science Advances, 7, eabd8800, 10.1126/sciadv.abd8800, 2021.

Berger, C. J. M., Lippiatt, S. M., Lawrence, M. G., and Bruland, K. W.: Application of a chemical leach technique for estimating labile particulate aluminum, iron, and manganese in the Columbia River plume and coastal waters off Oregon and Washington, J. Geophys. Res.-Oceans, 113, C00B01, https://doi.org/10.1029/2007JC004703, 2008.

Buck, C. S., Landing, W. M., and Resing, J.: Pacific Ocean aerosols: Deposition and solubility of iron, aluminum, and other trace elements, Marine Chemistry, 157, 117-130, 2013.

Buck, C. S., Landing, W. M., Resing, J. A., and Lebon, G. T.: Aerosol iron and aluminum solubility in the northwest Pacific Ocean: Results from the 2002 IOC cruise, Geochem. Geophys. Geosyst., 7, Q04M07, doi: 10.1029/2005GC000977, 2006.

Buck, C. S., Landing, W. M., Resing, J. A., and Measures, C. I.: The solubility and deposition of aerosol Fe and other trace elements in the North Atlantic Ocean: Observations from

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

659 the A16N CLIVAR/CO2 repeat hydrography section, Mar. Chem., 120, 57-70, 2010.
660 Buck, C. S., Fietz, S., Hamilton, D. S., Ho, T. Y., Perron, M. M. G., and Shelley, R. U.:
661 GEOTRACES: Fifteen years of progress in marine aerosol research, Oceanography,
662 37, 116-119, 2024.
663 Carslaw, D. C. and Ropkins, K.: openair - An R package for air quality data analysis,
664 Environmental Modelling & Software, 27-28, 52-61, 2012.
665 Cassar, N., Bender, M. L., Barnett, B. A., Fan, S., Moxim, W. J., Levy, H., and Tilbrook, B.:
666 The Southern Ocean biological response to Aeolian iron deposition, Science, 317,
667 1067-1070, 2007.
668 Chen, Y., Street, J., and Paytan, A.: Comparison between pure-water- and seawater-soluble
669 nutrient concentrations of aerosols from the Gulf of Aqaba, Mar. Chem., 101, 141-152,
670 2006.
671 Chen, Y. Z., Wang, Z. Y., Fang, Z. Y., Huang, C. P., Xu, H., Zhang, H. H., Zhang, T. Y.,
672 Wang, F., Luo, L., Shi, G. L., Wang, X. M., and Tang, M. J.: Dominant Contribution
673 of Non-dust Primary Emissions and Secondary Processes to Dissolved Aerosol Iron,
674 Environ. Sci. Technol., 58, 17355-17363, 2024.
675 Clough, R., Lohan, M. C., Ussher, S. J., Nimmo, M., and Worsfold, P. J.: Uncertainty
676 associated with the leaching of aerosol filters for the determination of metals in aerosol
677 particulate matter using collision/reaction cell ICP-MS detection, Talanta, 199, 425-
678 430, 2019.
679 Fishwick, M. P., Sedwick, P. N., Lohan, M. C., Worsfold, P. J., Buck, K. N., Church, T. M.,
680 and Ussher, S. J.: The impact of changing surface ocean conditions on the dissolution
681 of aerosol iron, Glob. Biogeochem. Cycle, 28, 1235-1250,
682 https://doi.org/10.1002/2014GB004921, 2014.
683 Gao, Y., Yu, S., Sherrell, R. M., Fan, S., Bu, K., and Anderson, J. R.: Particle-Size Distributions
684 and Solubility of Aerosol Iron Over the Antarctic Peninsula During Austral Summer,
685 J. Geophys. Res.-Atmos, 125, e2019JD032082, doi: 032010.031029/032019jd032082,
686 2020.
687 Hamilton, D. S., Moore, J. K., Arneth, A., Bond, T. C., Carslaw, K. S., Hantson, S., Ito, A.,
688 Kaplan, J. O., Lindsay, K., Nieradzik, L., Rathod, S. D., Scanza, R. A., and Mahowald,
689 N. M.: Impact of Changes to the Atmospheric Soluble Iron Deposition Flux on Ocean
690 Biogeochemical Cycles in the Anthropocene, Glob. Biogeochem. Cycle, 34,
691 10.1029/2019gb006448, 2020.
692 Hsu, S.-C., Lin, F.-J., and Jeng, W.-L.: Seawater solubility of natural and anthropogenic metals
693 within ambient aerosols collected from Taiwan coastal sites, Atmos. Environ., 39,
694 3989-4001, 2005.
695 Ito, A., Ye, Y., Baldo, C., and Shi, Z. B.: Ocean fertilization by pyrogenic aerosol iron, NPJ
696 Clim. Atmos. Sci., 4, 30, doi: 10.1038/s41612-41021-00185-41618, 2021.
697 Jaccard, S. L., Hayes, C. T., Martínez-García, A., Hodell, D. A., Anderson, R. F., Sigman, D.
698 M., and Haug, G. H.: Two Modes of Change in Southern Ocean Productivity Over the
699 Past Million Years, Science, 339, 1419-1423, 2013.
700 Jiang, H., Tang, J., Li, J., Zhao, S., Mo, Y., Tian, C., Zhang, X., Jiang, B., Liao, Y., Chen, Y.,
701 and Zhang, G.: Molecular Signatures and Sources of Fluorescent Components in
702 Atmospheric Organic Matter in South China, Environ. Sci. Tech. Lett., 9, 913-920,
703 2022.
704 Jickells, T. D., Baker, A. R., and Chance, R.: Atmospheric transport of trace elements and
705 nutrients to the oceans, Philosophical Transactions of the Royal Society a-
706 Mathematical Physical and Engineering Sciences, 374, 20150286, doi:
707 20150210.20151098/rsta.20152015.20150286, 2016.
708 Jickells, T. D., An, Z. S., Andersen, K. K., Baker, A. R., Bergametti, G., Brooks, N., Cao, J. J.,

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

Boyd, P. W., Duce, R. A., Hunter, K. A., Kawahata, H., Kubilay, N., laRoche, J., Liss, P. S., Mahowald, N., Prospero, J. M., Ridgwell, A. J., Tegen, I., and Torres, R.: Global Iron Connections between Desert Dust, Ocean Biogeochemistry, and Climate, Science, 308, 67-71, 2005.

Jordi, A., Basterretxea, G., Tovar-Sánchez, A., Alastuey, A., and Querol, X.: Copper aerosols inhibit phytoplankton growth in the Mediterranean Sea, Proc. Natl. Acad. Sci. U. S. A., 109, 21246-21249, 2012.

Kanthale, P., Ashokkumar, M., and Grieser, F.: Sonoluminescence, sonochemistry (H_2O_2 yield) and bubble dynamics: Frequency and power effects, Ultrasonics Sonochemistry, 15, 143-150, 2008.

Kumar, A., Sarin, M. M., and Srinivas, B.: Aerosol iron solubility over Bay of Bengal: Role of anthropogenic sources and chemical processing, Mar. Chem., 121, 167-175, 2010.

Langmann, B., Zakšek, K., Hort, M., and Duggen, S.: Volcanic ash as fertiliser for the surface ocean, Atmos. Chem. Phys., 10, 3891-3899, 2010.

Li, R., Panda, P. P., Chen, Y. Z., Zhu, Z. M., Wang, F., Zhu, Y. J., Meng, H., Ren, Y., Kumar, A., and Tang, M. J.: Aerosol trace element solubility determined using ultrapure water batch leaching: an intercomparison study of four different leaching protocols, Atmos. Meas. Tech., 17, 3147-3156, 2024.

Li, R., Dong, S. W., Huang, C. P., Yu, F., Wang, F., Li, X. F., Zhang, H. H., Ren, Y., Guo, M. X., Chen, Q. C., Ge, B. Z., and Tang, M. J.: Evaluating the effects of contact time and leaching solution on measured solubilities of aerosol trace metals, Appl. Geochem., 148, 105551, https://doi.org/10.1016/j.apgeochem.2022.105551, 2023.

Lu, L., Li, L., Rathod, S., Hess, P., Martínez, C., Fernandez, N., Goodale, C., Thies, J., Wong, M. Y., Alaimo, M. G., Artaxo, P., Barraza, F., Barreto, A., Beddows, D., Chellam, S., Chen, Y., Chuang, P., Cohen, D. D., Dongarrà, G., Gaston, C., Gómez, D., Morera-Gómez, Y., Hakola, H., Hand, J., Harrison, R., Hopke, P., Hueglin, C., Kuang, Y.-W., Kyllönen, K., Lambert, F., Maenhaut, W., Martin, R., Paytan, A., Prospero, J., González, Y., Rodriguez, S., Smichowski, P., Varrica, D., Walsh, B., Weagle, C., Xiao, Y.-H., and Mahowald, N.: Characterizing the Atmospheric Mn Cycle and Its Impact on Terrestrial Biogeochemistry, Glob. Biogeochem. Cycle, 38, e2023GB007967, https://doi.org/10.1029/2023GB007967, 2024.

Mackey, K. R. M., Chien, C.-T., Post, A. F., Saito, M. A., and Paytan, A.: Rapid and gradual modes of aerosol trace metal dissolution in seawater, Frontiers in Microbiology, 5, 794, 10.3389/fmicb.2014.00794, 2015.

Mahowald, N. M., Hamilton, D. S., Mackey, K. R. M., Moore, J. K., Baker, A. R., Scanza, R. A., and Zhang, Y.: Aerosol trace metal leaching and impacts on marine microorganisms, Nature Comm., 9, 2614, 2018.

Meskhidze, N., Volker, C., Al-Abadleh, H. A., Barbeau, K., Bressac, M., Buck, C., Bundy, R. M., Croot, P., Feng, Y., Ito, A., Johansen, A. M., Landing, W. M., Mao, J. Q., Myriokefalitakis, S., Ohnemus, D., Pasquier, B., and Ye, Y.: Perspective on identifying and characterizing the processes controlling iron speciation and residence time at the atmosphere-ocean interface, Mar. Chem., 217, 103704, 2019.

Miljevic, B., Hedayat, F., Stevanovic, S., Fairfull-Smith, K. E., Bottle, S. E., and Ristovski, Z. D.: To Sonicate or Not to Sonicate PM Filters: Reactive Oxygen Species Generation Upon Ultrasonic Irradiation, Aerosol Science and Technology, 48, 1276-1284, 2014.

Miller, J. N., and Miller, J. C.: Statistics and Chemometrics for Analytical Chemistry, 6th Edition, 2010.

Morton, P. L., Landing, W. M., Hsu, S.-C., Milne, A., Aguilar-Islas, A. M., Baker, A. R., Bowie, A. R., Buck, C. S., Gao, Y., Gichuki, S., Hastings, M. G., Hatta, M., Johansen, A. M., Losno, R., Mead, C., Patey, M. D., Swarr, G., Vandermark, A., and Zamora, L.

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

M.: Methods for the sampling and analysis of marine aerosols: results from the 2008 GEOTRACES aerosol intercalibration experiment, Limnol. Oceanogr. Methods, 11, 62-78, 2013.

Panda, P. P., Aswini, M. A., Bhatt, P., Srimuruganandam, B., Peketi, A., and Kumar, A.: Bioactive Trace Elements' Composition and Their Fractional Solubility in Aerosols from the Arabian Sea during the Southwest Monsoon, ACS Earth and Space Chem., 6, 1969-1981, 2022.

Paris, R. and Desboeufs, K. V.: Effect of atmospheric organic complexation on iron-bearing dust solubility, Atmos. Chem. Phys., 13, 4895-4905, 2013.

Paytan, A., Mackey, K. R. M., Chen, Y., Lima, I. D., Doney, S. C., Mahowald, N., Labiosa, R., and Post, A. F.: Toxicity of atmospheric aerosols on marine phytoplankton, Proc. Natl. Acad. Sci. U.S.A., 106, 4601-4605, 2009.

Perron, M. M. G., Fietz, S., Hamilton, D. S., Ito, A., Shelley, R. U., and Tang, M. J.: Preface to the inter-journal special issue "RUSTED: Reducing Uncertainty in Soluble aerosol Trace Element Deposition", Atmos. Meas. Tech., 17, 165-166, 2024.

Perron, M. M. G., Strzelec, M., Gault-Ringold, M., Proernse, B. C., Boyd, P. W., and Bowie, A. R.: Assessment of leaching protocols to determine the solubility of trace metals in aerosols, Talanta, 208, 120377, DOI: 120310.121016/j.talanta.122019.120377, 2020a.

Perron, M. M. G., Proemse, B. C., Strzelec, M., Gault-Ringold, M., Boyd, P. W., Rodriguez, E. S., Paull, B., and Bowie, A. R.: Origin, transport and deposition of aerosol iron to Australian coastal waters, Atmos. Environ., 228, 117432, 10.1016/j.atmosenv.2020.117432, 2020b.

Sarthou, G., Baker, A. R., Blain, S., Achterberg, E. P., Boye, M., Bowie, A. R., Croot, P., Laan, P., de Baar, H. J. W., Jickells, T. D., and Worsfold, P. J.: Atmospheric iron deposition and sea-surface dissolved iron concentrations in the eastern Atlantic Ocean, Deep Sea Research Part I: Oceanographic Research Papers, 50, 1339-1352, 2003.

Sedwick, P. N., Sholkovitz, E. R., and Church, T. M.: Impact of anthropogenic combustion emissions on the fractional solubility of aerosol iron: Evidence from the Sargasso Sea, Geochem. Geophys. Geosyst., 8, Q10Q06, DOI: 10.1029/2007GC001586, 2007.

Shelley, R. U., Perron, M. G. G., Hamilton, D. S., and Ito, A.: The open ocean, aerosols, and every other breath you take, Eos, 105, https://doi.org/10.1029/2024EO240091, 2024.

Shi, Z. B., Krom, M. D., Jickells, T. D., Bonneville, S., Carslaw, K. S., Mihalopoulos, N., Baker, A. R., and Benning, L. G.: Impacts on iron solubility in the mineral dust by processes in the source region and the atmosphere: A review, Aeolian Res., 5, 21-42, 2012.

Sholkovitz, E. R., Sedwick, P. N., Church, T. M., Baker, A. R., and Powell, C. F.: Fractional solubility of aerosol iron: Synthesis of a global-scale data set, Geochim. Cosmochim. Acta, 89, 173-189, 2012.

Stein, A. F., Draxler, R. R., Rolph, G. D., Stunder, B. J. B., Cohen, M. D., and Ngan, F.: NOAA's HYSPLIT Atmospheric Transport and Dispersion Modeling System, Bull. Amer. Meteorol. Soc., 96, 2059-2077, 2015.

Tang, W., Llort, J., Weis, J., Perron, M. M. G., Basart, S., Li, Z., Sathyendranath, S., Jackson, T., Sanz Rodriguez, E., Proemse, B. C., Bowie, A. R., Schallenberg, C., Strutton, P. G., Matear, R., and Cassar, N.: Widespread phytoplankton blooms triggered by 2019–2020 Australian wildfires, Nature, 597, 370-375, 2021.

Winton, V. H. L., Bowie, A. R., Edwards, R., Keywood, M., Townsend, A. T., van der Merwe, P., and Bollhofer, A.: Fractional iron solubility of atmospheric iron inputs to the Southern Ocean, Marine Chemistry, 177, 20-32, 2015.

Yu, X., Pan, Y., Song, W., Li, S., Li, D., Zhu, M., Zhou, H., Zhang, Y., Li, D., Yu, J., Wang, X., and Wang, X.: Wet and Dry Nitrogen Depositions in the Pearl River Delta, South

https://doi.org/10.5194/egusphere-2025-3274
Preprint. Discussion started: 17 July 2025

China: Observations at Three Typical Sites With an Emphasis on Water-Soluble Organic Nitrogen, J. Geophys. Res.-Atmos, 125, e2019JD030983, https://doi.org/10.1029/2019JD030983, 2020.

Zhang, H. H., Li, R., Dong, S. W., Wang, F., Zhu, Y. J., Meng, H., Huang, C. P., Ren, Y., Wang, X. F., Hu, X. D., Li, T. T., Peng, C., Zhang, G. H., Xue, L. K., Wang, X. M., and Tang, M. J.: Abundance and Fractional Solubility of Aerosol Iron During Winter at a Coastal City in Northern China: Similarities and Contrasts Between Fine and Coarse Particles, J. Geophys. Res.-Atmos, 127, e2021JD036070, DOI: 036010.031029/032021JD036070, 2022.