



1 **Impact Study of Increased Radio Occultation Observations during the** 2 **ROMEX Period Using JEDI and the GFS Atmospheric Model**

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8 **Abstract.** The international collaborative Radio Occultation Modeling EXperiment (ROMEX)
9 project marks the first time using a large volume of real data to assess the impact of increased
10 Global Navigation Satellite System (GNSS) radio occultation (RO) observations beyond
11 current operational levels, moving past previous theoretical simulation-based studies. The
12 ROMEX project enabled the use of approximately 35,000 RO profiles– nearly triple the
13 number typically available to operational centers, which is about 8,000 to 12,000 per day. This
14 study investigates the impact of increased RO profiles on numerical weather prediction (NWP)
15 with the Joint Effort for Data assimilation Integration (JEDI) and the global forecast system
16 (GFS), as part of the ROMEX effort. A series of experiments were conducted assimilating
17 varying amounts of RO data along with a common set of other key observations. The results
18 confirm that assimilating additional RO data further improves forecasts across all major
19 meteorological fields, including temperature, humidity, geopotential height, and wind speed,
20 for most of vertical levels. These improvements are significantly evident in verification against
21 both critical observations and the European Center for Medium-Range Weather Forecasts
22 (ECMWF) analyses, with beneficial impacts lasting up to five days. Conversely, withholding
23 RO data resulted in forecast degradations. The results also suggest that forecast improvements
24 scale approximately logarithmically with the number of assimilated profiles, and no evidence
25 of saturation was observed. Biases in the forecast of temperature and geopotential height over
26 the lower stratosphere are discussed, and they are consistent with findings from other studies
27 in the ROMEX community.

28



29 **1 Introduction**

30 Radio occultation (RO) is an active remote sensing technique that measures the refraction
31 of signals transmitted by Global Navigation Satellite System (GNSS) and received by
32 instruments aboard low-Earth orbit (LEO) satellites. The pioneering GPS/Meteorology
33 (GPS/MET) mission demonstrated that the RO technique can effectively probe the Earth's
34 atmosphere, providing profiles with high vertical resolution and accuracy (Kursinski et al.,
35 1997). Since then, the number of RO profiles has increased with the expansion of GNSS
36 beyond GPS (e.g. GLONASS, Galileo, and BeiDou), along with the deployment of more RO
37 receivers aboard new LEO missions, such as the U.S./Taiwan FORMOSAT-3/Constellation
38 Observing System for Meteorology, Ionosphere, and Climate (COSMIC-1; launched in 2006)
39 and its successor FORMOSAT-7/COSMIC-2 (COSMIC-2, launched in 2019), the European
40 Space Agency (ESA) MetOp series (MetOp-A, 2006; MetOp-B, 2012; MetOp-C, 2018), and
41 the ESA/US Sentinel-6 (launched in 2020).

42 RO data are considered as one of the most impactful observation types in terms of their
43 contribution to the forecast skills in numerical weather prediction (NWP). The positive impact
44 of RO observations on NWP analysis and forecast has been well-documented by numerous
45 NWP centers (Healy and Thépaut, 2006; Bowler 2020; Ruston and Healy 2021; Cucurull
46 2023). Unlike satellite radiance data, RO observations are inherently unaffected by clouds or
47 precipitation and therefore their assimilation in NWP requires no bias correction. Instead, RO
48 observations serve as a reference to anchor the bias correction of satellite radiance data (Healy
49 et al. 2008; Bauer et al. 2014).

50 Since launched in 2019, COSMIC-2 has been steadily providing approximately 6,000 RO
51 profiles daily, primarily between 45°S and 45°N. Other government missions in polar orbits
52 contribute around 2,000 daily profiles globally. More recently, the emergence of commercial
53 RO providers, such as GeoOptics Inc. (Pasadena, CA, USA), PlanetiQ (Golden, CO, USA) and
54 Spire Global Inc. (Boulder, CO, USA), has further expanded RO data availability. These
55 commercial sources supplement operational capabilities with additional profiles, depending on
56 purchase agreements. With this expanded data volume, NWP centers have recently been
57 enabled to explore the impact of assimilating slightly more than 10,000 RO profiles per day.



58 Several NWP centers have demonstrated that the relative impact of RO data in NWP has grown
59 alongside the increasing availability of profiles. Bowler (2020) at the Met Office assessed the
60 RO data produced by Spire and stated that the forecast impact of increasing the RO data volume
61 is roughly proportional to the logarithm of the total amount of GNSS RO data assimilated.
62 Ruston and Healy (2021) reported a novel finding that COSMIC-2 data improve the tropical
63 tropospheric humidity forecasts in both the Navy Global Environmental Model (NAVGEM)
64 and the European Center for Medium-Range Weather Forecasts (ECMWF) Integrated
65 Forecasting System (IFS). The assimilation of COSMIC-2 and Spire observations was found
66 beneficially in both the ECMWF and Met Office's NWP systems (Lonitz et al. 2021). Cucurull
67 (2023) demonstrated that COSMIC-2 observations have a significant impact in the forecast
68 improvement of temperature and winds in the tropics.

69 While these studies demonstrated the valuable impact of increased RO profiles in
70 operational NWP systems, the potential benefits of even larger data volumes were only
71 explored through theoretical simulations. Harnisch et al. (2013) used an ensemble of data
72 assimilations (EDA) approach to evaluate the change of RO data impact as a function of
73 observation numbers. They demonstrated that saturation was not found with 128,000 simulated
74 RO profiles per day. With a global observing system simulation experiment (OSSE) study,
75 Privé et al. (2022) found that the assimilation of 100,000 daily RO profiles did not reach the
76 impact saturation in the hybrid four-dimensional ensemble variational data assimilation system
77 and Global Earth Observing System (GEOS) model.

78 Clearly, the number of RO observations currently utilized in real-time NWP operations
79 remains significantly below the potential demonstrated in these simulated studies. Meanwhile,
80 a large portion of RO observations from commercial providers is not purchased and remains
81 unassimilated in operational systems, highlighting that the full impact of RO from both
82 government and commercial providers has yet to be fully realized. Since May 2022, the
83 International RO Working Group (IROWG; <https://irowg.org>), one of the scientific advisory
84 working groups of the Coordination Group for Meteorological Satellites (CGMS), has led an
85 international collaborative effort – the Radio Occultation Modeling Experiment (ROMEX;
86 Anthes et al. 2024) – to explore the impact of RO observations by collecting as many profiles
87 as available from both commercial and government providers during the testing period.



88 Specifically, ROMEX has collected nearly 35,000 daily profiles during the experimental period
89 (September to November 2022), whereas there are about 12,000 daily profiles available in the
90 real-time NWP operations at the U.S. National Oceanic and Atmospheric Administration
91 (NOAA). ROMEX provides a unique opportunity for both NWP centers and the research
92 community to evaluate impacts of increased RO numbers using large quantities of real RO
93 observations for the first time.

94 The overarching objective of this study is to demonstrate forecast improvement through the
95 assimilation of the increased RO data. We aim to quantitatively assess these data impacts with
96 respect to operational implementations, while leveraging advanced features for enhanced
97 performance. Specifically, this study utilizes the Joint Effort for Data assimilation Integration
98 (JEDI; Trémolet and Auligné 2020) for data assimilation and the NOAA Global Forecast
99 System (GFS) for forecasting. Given JEDI is the next generation data assimilation system for
100 operations at NOAA, the National Aeronautics and Space Administration (NASA), the U.S.
101 Naval Research Laboratory (NRL), and other NWP centers worldwide, this ROMEX study
102 offers additional benefits by demonstrating JEDI's capabilities and providing insights for
103 ongoing transitions to operations.

104 This manuscript is organized as follows: Section 2 summarizes ROMEX and the RO
105 observations used for this study; Section 3 introduces the GFS forecast model, the JEDI data
106 assimilation system, and the experimental design; Section 4 presents the evaluation of the
107 ROMEX RO data impact using the JEDI-GFS system; Section 5 presents a summary of the
108 work.

109 **2 ROMEX and GNSS RO observations**

110 ROMEX is an IROWG initiative designed to evaluate the impact of increasing radio
111 occultation (RO) data volume using real observations from both government and commercial
112 missions, extending beyond current operational capabilities. ROMEX involves approximately
113 30 international agencies and research institutions, including data providers, processing
114 centers, NWP centers, universities, and research institutes. A complete list of ROMEX
115 participants is available on the ROMEX website (<https://irowg.org/ro-modeling-experiment->



116 romex/). The outcomes of ROMEX provide guidance to CGMS for formulating
117 recommendations to space agencies on RO mission planning and coordination. Additionally,
118 ROMEX results offer valuable insights for RO data providers, processing centers, and NWP
119 centers to enhance data retrieval techniques and improve the assimilation of RO data in weather
120 forecasting.

121 Through dedicated data agreements with commercial RO providers, the ROMEX effort
122 was able to access data not covered by existing global licenses held by NOAA, NASA, or
123 EUMETSAT from their respective commercial purchases. The ROMEX dataset includes
124 commercial RO data otherwise unavailable to the public, along with publicly available data
125 from sources such as the UCAR COSMIC Data Acquisition and Access Center (CDAAC),
126 NOAA, NASA, and EUMETSAT. Due to the involvement of multiple processing centers and
127 data providers, different processing versions of the data were available to support validation
128 and processing studies. For this study, ROMEX data were processed either by UCAR,
129 EUMETSAT, or their original providers. All these data are distributed through the
130 EUMETSAT Radio Occultation Meteorology Satellite Application Facility (ROM SAF).

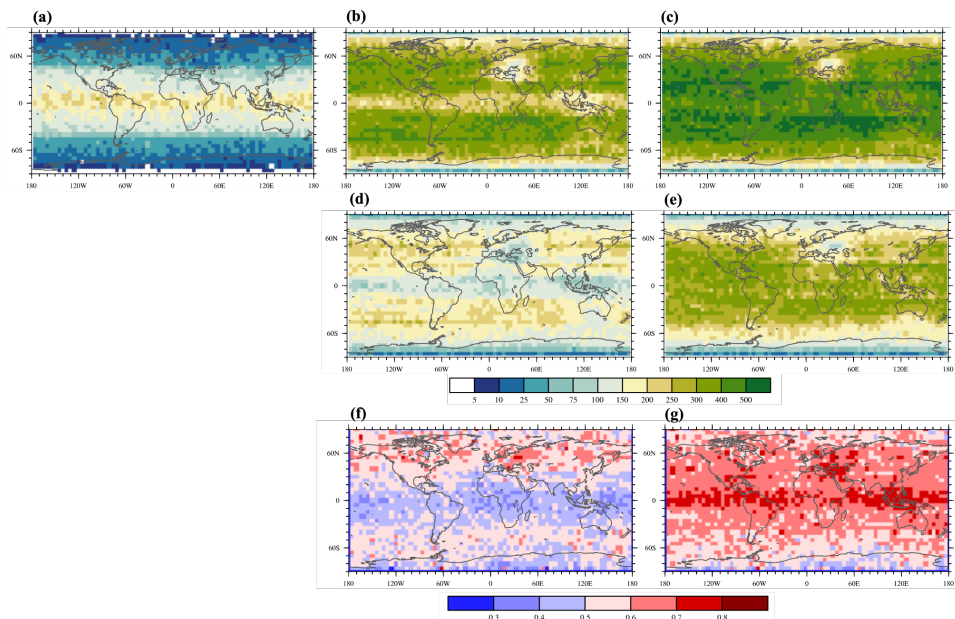
131 Our objective is to evaluate the impact of an increased volume of RO profiles on analyses
132 and forecasts, rather than comparing the performance or characteristics of various missions.
133 Early data evaluation already shows the quality of these data is relatively comparable for NWP
134 applications (Marquardt 2024; Anthes et al. 2025). The available profiles are categorized into
135 two groups based on their sources: base missions (hereafter, BASE) which are government-
136 supported missions, and the supplementary missions that are provided by commercial vendors.
137 The base missions include COSMIC-2, Metop-B, Metop-C, Kompsat-5, PAZ, Sentinel-6,
138 TerraSAR-X, and TanDEM-X, all of which are available in near-real-time for operational
139 NWP systems. The supplementary missions consist of GeoOptics, PlanetiQ (Kursinski 2025),
140 Spire (Nguyen 2025), Yunyao (Cheng 2025), and Tianmu (Tang 2025). On average,
141 approximately 35,000 profiles (8,000 from base missions and 27,000 from supplementary
142 missions) were available during the ROMEX period (hereafter, ROMEX). To further quantify
143 the impact of the increased profile volume, the EUMETSAT ROM SAF provided a sub-dataset
144 referred to as ROMEX20K, in which the average daily number of profiles is 20,000. In the



145 ROMEX20K sub-dataset, the supplementary profiles are reduced to approximately 12,000 per
146 day.

147 Figure 1 presents the total number of RO profiles in each $5^\circ \times 5^\circ$ latitude-longitude grid for
148 September 2022, the testing period of this study. Specifically, Fig. 1a–c shows the number of
149 BASE profiles, supplementary mission profiles, and all available ROMEX profiles,
150 respectively. Fig. 1d–e displays the supplementary profiles used in the ROMEX20K and all
151 ROMEX20K profiles. Fig. 1f–g shows the ratio of supplementary profiles used in ROMEX20K
152 with all supplementary profiles, and the ROMEX20K profiles relative to the total number of
153 ROMEX profiles, respectively.

154 The total number of BASE profiles (Fig. 1a) peaks in the tropics and decreases poleward,
155 primarily due to the dominance of COSMIC-2. The supplementary profiles (Fig. 1b), however,
156 are more evenly distributed across the mid-to-high latitudes. Overall, the combination of all
157 available profiles (Fig. 1c) results in a relatively uniform global distribution geographically.
158 The supplementary profiles used in the ROMEX20K sub-dataset kept a higher portion over the
159 northern hemisphere and southern mid-to-high latitudes than over the tropical regions (Fig. 1f).
160 Combined with the base profiles, ROMEX20K has a better coverage over the tropics than other
161 regions of the globe and the fewest profiles over the southern polar regions (Fig. 1g).



162

163 **Figure 1:** Number of RO profiles in September 2022, re-gridded to a $5^\circ \times 5^\circ$ latitude–longitude grid,
164 for the following data sets: (a) base missions (BASE), (b) all supplementary missions, (c) all ROMEX
165 missions (ROMEX), (d) supplementary missions in the ROMEX20K, (e) all missions used in the
166 ROMEX20K configuration. Panels (f) and (g) show the ratio of total profile counts between (d)
167 supplementary in ROMEX20K and (b) total supplementary missions, and between (e) ROMEX20K
168 and (c) all ROMEX missions, respectively.
169

170 **3 Forecast model, data assimilation, and experimental design**

171 **3.1 Forecast model**

172 The Global Forecast System (GFS) is NOAA’s medium-range operational global weather
173 prediction model, developed and maintained by the National Centers for Environmental
174 Prediction (NCEP). It is part of the Unified Forecast System (UFS), a community-based,
175 coupled Earth modeling framework designed to integrate research and operational weather
176 modeling for more consistent and advanced forecasts (Zhou et al. 2022). This study used the
177 atmospheric forecast model component of UFS, and not the entire suite. Further, the next



178 planned release version 17 implementation¹ (GFSv17; GFS hereafter) was employed in this
179 study. This latest version of GFS is continuing to use the Finite-Volume Cubed-Sphere (FV3;
180 Lin 2004) dynamical core, and also incorporates significant upgrades in parameterizations for
181 atmospheric processes such as cloud microphysics (Stefanova et al. 2022; Meixner et al. 2023),
182 in comparison to the current operational implementation at NCEP. The global forecasts for this
183 study are configured at a horizontal resolution of approximately 25 km with 127 vertical levels
184 extending up to 80 km (C384L128). This is half of the operational resolution and is standard
185 practice for pre-implementation testing at NCEP and by associated researchers.

186 3.2 Data assimilation

187 This study uses the Joint Effort for Data assimilation Integration (JEDI; Trémolet and
188 Auligné 2020) to fulfill the data assimilation component. Led by the Joint Center for Satellite
189 Data Assimilation (JCSDA), JEDI was initiated in 2017. As the project has grown partners
190 now include NOAA, NASA, NRL, the U.S. Air Force, the NSF National Center for
191 Atmospheric Research (NCAR), UK Met Office and developers from universities. JEDI has
192 been interfaced to various models, including the GFS, through the JEDI-FV3 component
193 (<https://github.com/JCSDA/fv3-jedi/>), allowing the partner agencies using FV3 core-based
194 systems to implement JEDI in real-time applications.

195 The observation operators for JEDI are developed within an abstracted and generic coding
196 layer known as the Unified Forward Operator (UFO). A generic design makes UFO model-
197 agnostic and allows it to be used in a play-and-plug manner through configuration files.
198 Currently, UFO includes six GNSS RO operators, four for bending angle and two for
199 refractivity, contributed by different partners to replicate the implementation in their respective
200 NWP systems. The four bending angle operators include one based on the operational NCEP
201 Bending Angle Model (NBAM; Cucurull et al. 2013), the Met Office's bending angle operator
202 (Burrows 2014, Burrows et al. 2014), and both one-dimensional (ROPP1D; Healy and Thépaut

¹GFSv17 has not been implemented in the operation as we started this work to the best of our knowledge. We checked out the branch prototype/hr3 in August 2024 from <https://github.com/ufs-community/ufs-weather-model>



203 2006) and two-dimensional (ROPP2D; Healy et al. 2007; Healy 2014) operators interfaced via
204 the ROM-SAF Radio Occultation Processing Package (ROPP; [https://rom-](https://rom-saf.eumetsat.int/ropp)
205 [saf.eumetsat.int/ropp](https://rom-saf.eumetsat.int/ropp)), which are used operationally by NRL and ECMWF, respectively.
206 Additionally, two refractivity operators are included, following implementations from the Met
207 Office and NCEP (Cucurull et al. 2007, Buontempo et al. 2008). Most NWP centers use 1D
208 bending angle operators operationally, considering both its impacts and computational
209 efficiency. While a detailed comparison of these operators is performed in a separate effort, we
210 use the ROPP1D operator with the default JEDI configuration that was based on the current
211 implementations by partner agencies. UFO also includes associated quality control (QC)
212 procedures and observation error models, allowing creation of a consistent treatment to those
213 used in operational applications at other centers.

214 3.2.1 GNSS RO forward operator

215 Assuming the atmosphere is horizontally homogeneous and spherically symmetric,
216 ROPP1D computes the bending angle, α , by vertically integrating the refractive index from the
217 model background, as shown in Equation 1 (Healy and Thepaut 2006). a is the observed impact
218 parameter, n is the modelled refractive index, and $x = nr$ is the product of the refractive index
219 and the radius value r of a point on the ray path.

$$220 \quad \alpha(a) = -2a \int_a^\infty \frac{1}{\sqrt{x^2 - a^2}} \frac{d \ln(n)}{dx} dx \quad (1)$$

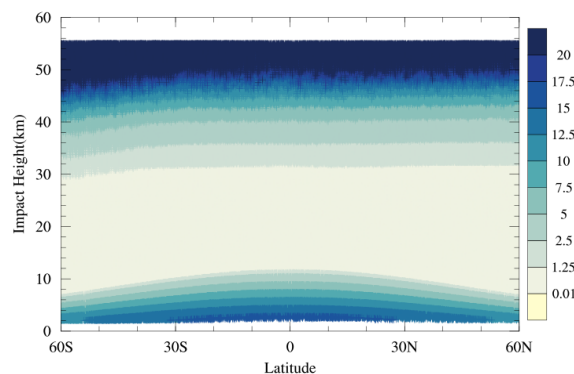
221 The model background information is extracted for each observation point along the RO
222 profile, valid at the horizontal location of its corresponding tangent point. Therefore, the
223 vertical drift of tangent points is fully accounted for. Impact height – defined as the difference
224 between the impact parameter and the local radius of curvature of the Earth – is referred to as
225 the vertical coordinate when presenting RO space results. Note that the impact parameter is a
226 geometric quantity representing the closest distance between the straight-line trajectory of a
227 GNSS signal and the Earth's center; it is the actual coordinate in RO assimilation.

228 3.2.2 RO observation error and quality control



229 Observation error and quality control procedure are two crucial parameters in DA. The
230 observation error accounts for measurement uncertainty, representativeness error, and forward
231 operator error (Bormann 2015). Accurate modeling of observation errors is essential for
232 appropriately weighting RO observations relative to other data types and the background error.
233 Meanwhile, QC procedures are closely linked to both the forward operator and observation
234 uncertainty, as they aim to remove observations with large departures that may result from
235 forward operator limitations or various sources of measurement error. As such, observation
236 error characterization, quality control, and the forward operator are tightly interconnected in
237 the assimilation of RO observations.

238 We applied the observation error model used in the NRL designed system (Ruston and
239 Healy 2021) that is run operationally at Fleet Numerical Meteorology and Oceanography
240 Center (FNMOC). Figure 2 shows the observation errors (in percentage) averaged over a
241 random day, as functions of latitude and impact height. In this scheme, observation errors are
242 defined as a percentage of the observed values and decrease linearly with increasing impact
243 height, reaching a minimum of 1.25 % at the “minimum error height”. A damping factor is
244 applied to account for latitudinal variation. In JEDI implementation, the error is specified as
245 20 % at the surface (impact height is 0) at 0° latitude and is reduced away from the equator
246 following the cosine of latitude. The minimum error height also varies with latitude, decreasing
247 from 12 km at the equator to 5.333 km at the poles – again modulated by the cosine of latitude.



248

249 **Figure 2:** Percentage observation errors (%) for all RO observations on September 1, 2022, using the
250 NRL error model.



251 **3.3 Experimental design**

252 Four sets of experiments—namely, noRO, BASE, ROMEX20K, and ROMEX—were
253 conducted using the JEDI-GFS system over a one-month period in September 2022. All
254 experiments assimilated a common set of conventional observations, cloud-motion vectors, and
255 satellite radiances from the Global Data Assimilation System (GDAS) archive. Specifically,
256 the observations including conventional data from radiosondes, aircraft, and surface stations,
257 as well as scatterometer wind vectors and radiances from AMSU-A and ATMS measurements
258 from multiple satellites available during the study period.

259 These four experiments differ only in the volume of RO profiles assimilated. The noRO
260 experiment excludes RO data entirely. The BASE experiment assimilates only publicly
261 available RO profiles, totaling approximately 8,000 per day. ROMEX20K and ROMEX
262 assimilate approximately 20,000 and 35,000 daily RO profiles, respectively, based on the
263 corresponding datasets. Differences in forecast skill among these experiments illustrate the
264 impact of enhanced RO data volume available during the ROMEX period.

265 All experiments were initialized at 0000 UTC on September 1, 2022, using a 6-hour
266 forecast as background from the operational GFS system at NCEP. The JEDI-GFS system was
267 then cycled every 6 hours, with a 6-hour assimilation window and background fields provided
268 by the previous forecast cycle. Data assimilation was performed using JEDI's hybrid three-
269 dimensional variational (3DVar) method, with 40 ensemble members taken from the NCEP
270 global ensemble forecast system. The data assimilation minimization was performed on a so-
271 called dual-resolution grid: the background and forecasts used the C384 grid, while the
272 minimization was carried out on the C192 grid.

273 **4 Results and evaluation**

274 This section compares the results of all experiments to assess the impact of RO observations
275 on forecast skill. Short-to-medium range forecasts are evaluated against observations and
276 model analyses. In observation space, common evaluation metrics include observation–minus–
277 background (OMB) and observation–minus–analysis (OMA) statistics, whose mean values are



often referred to as background bias or analysis bias, respectively. In model space, forecast skill is assessed by comparing model forecasts and ECMWF analyses (FMA) at analysis grid points. Three basic metrics, root-mean-square error (RMSE), standard deviation (STDV), and mean bias, are calculated for OMB/OMA or FMA over the entire experimental period.

To further evaluate the impact of each experiment relative to the reference (noRO or BASE), two additional metrics are adopted to illustrate the impact of an experiment relative to the reference experiment. The first is the normalized difference of a given metric between the experiment and the chosen reference (Eq. 2), where a negative value indicates improvement and a positive value indicates degradation. The second is the mean absolute error reduction (MAER; Eq. 3), which compares the absolute biases between experiments. A negative MAER reflects a beneficial bias reduction relative to the reference experiment, while a positive value indicates a detrimental increase.

$$\text{Normalized difference} = 100\% \times \frac{M(\text{Exp.}) - M(\text{Reference})}{M(\text{Reference})} \quad (2)$$

$$\text{MAER} = 100\% \times \frac{|\text{Bias}(\text{Exp.})| - |\text{Bias}(\text{Reference})|}{|\text{Bias}(\text{Reference})|} \quad (3)$$

4.1 Evaluation in observation space

Statistics in RO observation space are first calculated to evaluate the performance of the JEDI-GFS system in assimilating the large volume of real RO data from ROMEX. Because RO bending angle observation values span a few orders of magnitudes vertically, the OMB statistics are presented in a normalized format, i.e., OMB/O (B is a 6-h forecast from the previous cycle). Figure 3 shows the RMSE and bias from the ROMEX experiment. RMSE results (Fig. 3a) show that the assimilation produces a reasonable agreement between the bending angle observations and both the 6-h forecast and analysis, with lower RMSE in the analysis than in the background. Biases are also notably reduced after assimilation especially between 5 and 12 km impact heights, when comparing OMA to OMB (Fig. 3b), demonstrating the assimilation's effectiveness in correcting background errors in this key region.

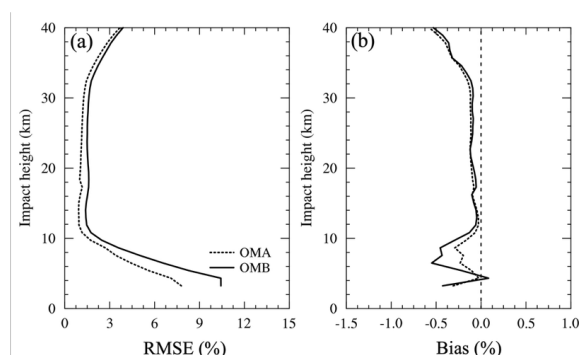


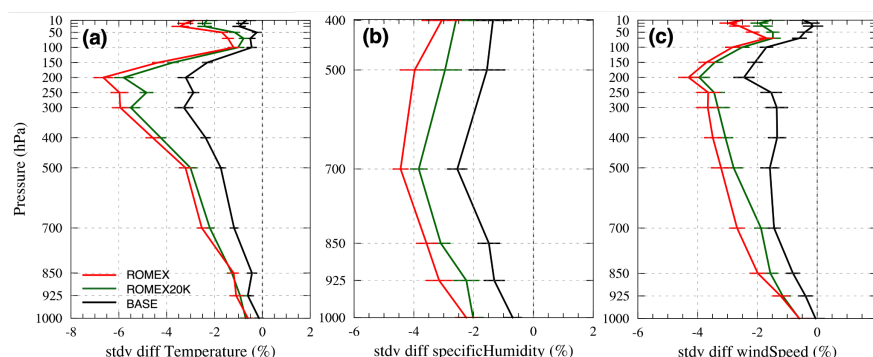
Figure 3: (a) RMSE and (b) Bias of OMB (RO bending angle observation minus 6-h model forecast) normalized by observations (OMB/O; solid) and OMA (RO bending angle observation minus analysis) similarly normalized (OMA/O; dashed), of the ROMEX experiment in September 2022.

Fig. 4 shows the normalized difference in STDV for the fit of temperature, specific humidity, and wind speed forecasts to radiosonde observations. All three experiments show smaller STDV than noRO across all variables and levels. The STDV reduction relative to noRO increases with height, reaching a maximum near 200 hPa for temperature and wind, and at around 700 hPa for humidity. Comparing all three RO experiments (ROMEX, ROMEX20K and BASE), the forecast improvements, or STDV reduction, increase with the growing volume of RO data. ROMEX generates the largest reduction among the three RO experiments. For example, the STDV reduction of temperature forecast at 200 hPa are 3.2%, 5.3%, and 6.8% for BASE, ROMEX20K, and ROMEX respectively. However, the difference between ROMEX and ROMEX20K is negligible in the lower troposphere (below 800 hPa) for temperature (Fig. 4a), and near the surface for wind speed (Fig. 4c). Note that RO data provide only information on the mass fields, wind forecasts are not directly impacted by the assimilation of RO data. Rather, they are impacted through the background error covariance between state variables in DA and the dynamic adjustment through the month-long cycles. Despite this indirect influence, the positive impact of RO data on wind forecasts seen here is very significant.

All RO experiments reduced noRO bias in temperature forecasts between 700 hPa and 100 hPa (Fig. 5). Experiments with additional RO data assimilation tend to cool nearly the entire troposphere, as indicated by the RO experiment curves lying to the right side of the NoRO



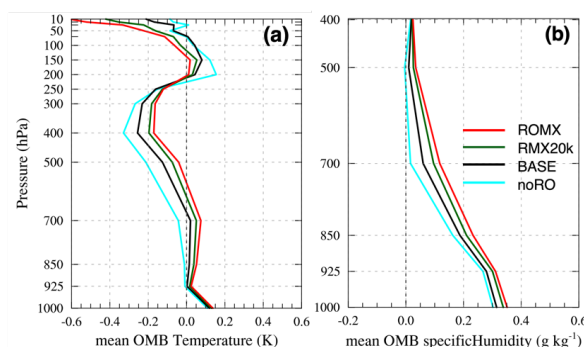
325 curve. In contrast, it warms the air above the jet-stream layer around 200 hPa, where the RO
326 experiment curves shift to the left of the NoRO curve (Fig. 5a). However, the RO experiments
327 produce larger humidity biases compared to noRO. Assimilating more RO data results in a
328 drier atmosphere, while withholding RO data leads to a wetter one (Fig. 5b).



329

330 **Figure 4:** Normalized difference in STDV (%) of the RO experiments relative to noRO for the 6 h model
331 forecasts verified against radiosonde observations of (a) temperature, (b) specific humidity, and (c)
332 wind speed. Overlaid bars are the standard deviations of the normalized STDV difference.

333



334

335 **Figure 5:** Bias of all experiments for the 6 h model forecasts verified against radiosonde observations
336 (OMB) of (a) temperature (unit: K) and (b) specific humidity (unit: g kg^{-1}).

337

338 The impact of assimilating ROMEX RO data on other critical observing systems is also
339 examined to understand whether the assimilation of RO data can indirectly enhance the
340 assimilation of other observation types. Figure 6 presents the normalized difference in STDV,



relative to noRO, for the fit to aircraft temperature and wind speed observations. Consistent forecast improvements are observed in both fields, aligning with the verification result from radiosonde data. Figure 7 presents the percentage difference in the number of assimilated observations relative to the BASE experiment for six observation types (radiosonde, aircraft, surface, scatterometer, satellite winds or atmospheric motion vectors, and GNSS RO observations), by comparing ROMEX20K, ROMEX, and noRO. BASE serves as the reference experiment to facilitate the inclusion of GNSS RO data in the comparison. Note also that the three RO experiments assimilate different RO datasets, which does not favor a direct comparison of the total number of assimilated RO observations across experiments. Therefore only the RO observations from the BASE dataset assimilated in all experiments are considered for the bars labeled GNSSRO in Fig. 7. The observations passing quality control in the OMB statistics are counted through the entire month for each type. Both ROMEX20K and ROMEX assimilate more data than BASE across all data types used in this study, whereas noRO assimilates clearly fewer. For example, ROMEX20K and ROMEX assimilate 0.88% and 0.59% more radiosonde observations than BASE, respectively, while noRO assimilates 0.73% less. This indicates that assimilating additional RO data brings the model analysis and short-range forecast closer to these observations and enables the use of more observational data. This approach clearly shows that both ROMEX and ROMEX20K increased the assimilation of BASE RO data compared to the BASE experiment.

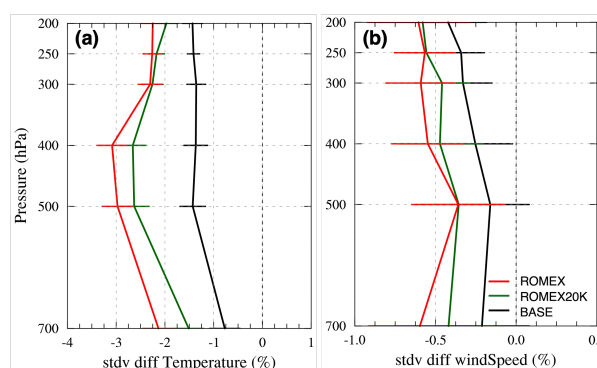
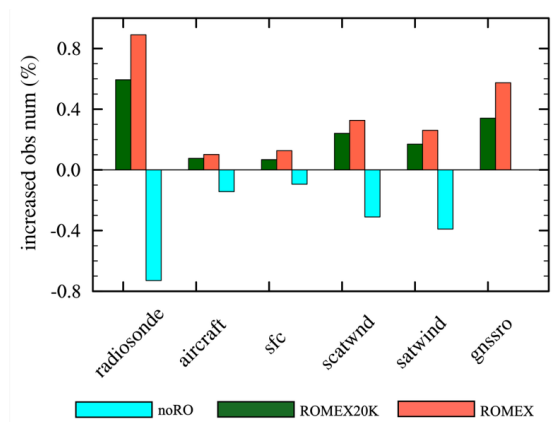


Figure 6: Fractional STDV difference (%) of experiments relative to noRO for the 6 h model forecasts verified against aircraft observations (OMB) of (a) temperature and (b) wind speed. Overlaid bars are the standard deviations of the fractional STDV difference.



365



366

367 **Figure 7:** Fractional observation count difference (%) of experiments relative to BASE for key
 368 observing types (radiosonde, aircraft, surface, scatterometer, satellite winds or atmospheric motion
 369 vectors, and GNSS RO observations from base missions) assimilated in all experiments. Note
 370 experiment BASE is used for reference here to account for the GNSS RO observation. Only the base
 371 GNSS RO mission observations are counted for the GNSS RO calculation.

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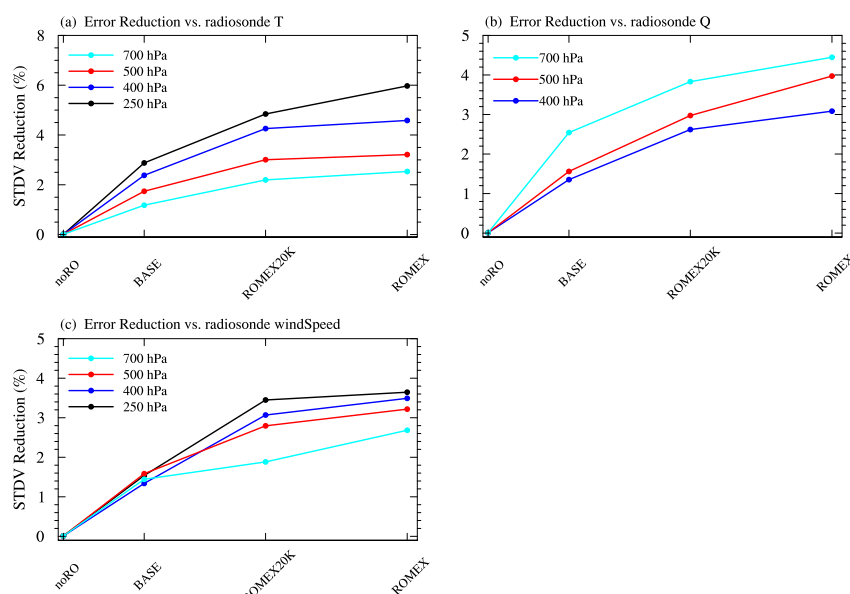


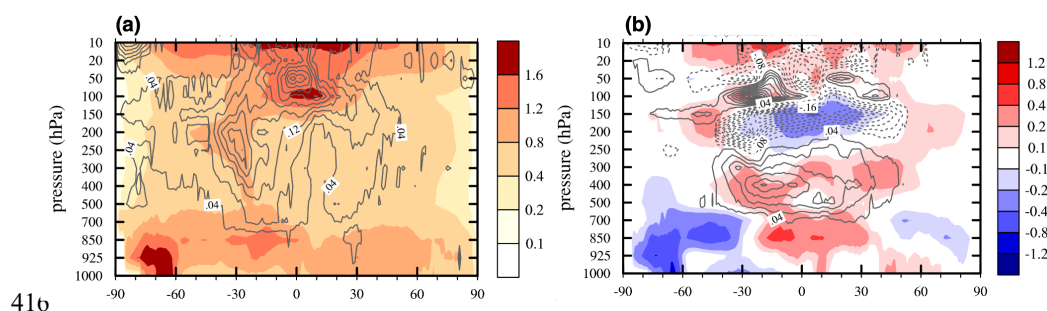
Figure 8: STDV reductions with the number of daily RO profiles for (a) temperature and (c) wind speed at 250-, 400-, 500-, and 700-hPa, and (b) specific humidity at 400-, 500-, and 700-hPa. X-axis is the experiments, and y-axis the fractional STDV difference relative to noRO. For illustration purposes, positive numbers are reductions for this figure.

To consolidate the STDV reductions for the forecast of temperature, specific humidity and wind at various levels, Figure 8 is presented to show the 6-h forecast STDV reductions for the various experiments that exhibit progressive increases in daily RO profiles at key pressure levels, following the NWP verification exchange guidance provided by ROMEX. Using noRO as the benchmark, the STDV reduction increases approximately logarithmically with the growing number of profiles, consistent with the findings of Lonitz (2025). Notably, there is no clear sign of saturation, as most levels continue to show improvement with increasing numbers of RO profiles. However, the degree of this non-saturation appears to depend on both the variable and vertical level and could be influenced by the specific data assimilation configuration.

4.2 Evaluation in model space



416 This sub-section evaluates the impact of ROMEX data on short-to-medium range forecasts
417 against the ECMWF analysis over the one-month experimental period. The first assessment
418 focuses on the noRO experiment by comparing its results with that of BASE, which represents
419 the consequences of losing or withholding all RO data. Figure 9 presents the zonal-mean of the
420 STDV and bias of the 6-h temperature forecasts of BASE as a function of pressure levels. Also
421 shown are the differences in STDV and bias between noRO and BASE, with statistical
422 significance at the 95% confidence level. The BASE experiment exhibits large temperature
423 STDVs in the tropical tropopause and lower stratosphere (above 50 hPa), as well as in the lower
424 troposphere, especially in the Southern Hemisphere high latitudes (Fig. 9a). When RO data are
425 withheld (noRO), significant forecast degradations (increasing STDV errors relative to the
426 ECMWF analysis) occur above 850 hPa, with a maximum increase in STDV, about 0.4 K,
427 centered near 50 hPa over the tropics. BASE's temperature bias exhibits multiple patterns,
428 including prominent negative values in the tropical tropopause, positive values above and
429 below the tropopause in the low latitudes, and negative values in the lower tropospheric at high
430 latitudes (Fig. 9b). noRO amplifies BASE's existing biases: negative values become more
431 negative, and positive values become more positive, particularly in the mid-to-upper tropical
432 troposphere and near the mid-latitude tropopause. The negative impact of excluding GNSS RO
433 observation as shown in noRO is also observed in the verification of other key parameters such
434 as humidity and wind speed (not shown), therefore the following sections will focus on the
435 presentation of the impact of the two ROMEX datasets relative to BASE.

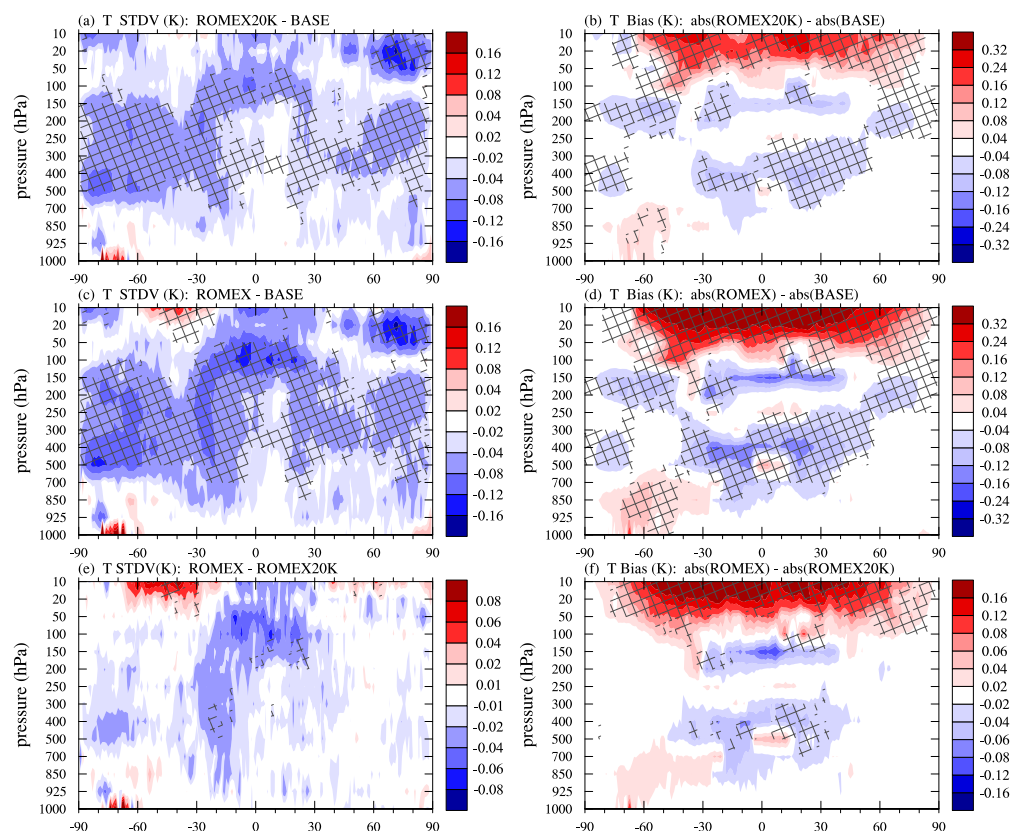


416 **Figure 9:** (a) Zonal-mean STDV (shaded), and (b) bias (shaded) of the 6-h temperature forecasts (unit:
417 K) of BASE verified against ECMWF analysis. Overlaid contours are the differences between noRO
418 and BASE (noRO-BASE; unit: K; interval: 0.04 K) in (a) STDV and (b) bias at 95% significance level.
419
420 Solid/dashed curves represent positive/negative values respectively in both panels. In panel (a), positive



values for contours indicate noRO forecast degradation relative to BASE, while negative values indicate improvement. In panel (b), opposite signs between the contours and shading indicate improvement in noRO relative to BASE, while matching signs indicate degradation.

Figure 10 compares the two ROMEX experiments, ROMEX and ROMEX20K, with respect to BASE in terms of STDV (Figs. 10a, c, and e) and MAER (Figs. 10b, d, and f) for 6-h temperature forecasts, and also includes a direct comparison between ROMEX and ROMEX20K. As introduced earlier, negative MAER values indicate reductions in absolute bias relative to BASE (i.e., improvement), while positive values indicate increased bias (i.e., degradation). Hashed areas indicate regions where the results are statistically significant at the 95% confidence level. Overall, both ROMEX and ROMEX20K exhibit significant STDV reductions relative to BASE between 850 and 50 hPa. Substantial forecast improvements are observed in the Southern Hemisphere's mid-to-upper troposphere, the tropical tropopause region, and the middle troposphere and stratosphere of the Northern Hemisphere's high latitudes. One exception is that ROMEX produces increased STDVs above 50 hPa over the Southern Hemisphere middle latitudes (Figs. 10a and c), which is linked to the warming effect introduced by the additional ROMEX RO data as seen in the verification against radiosonde temperatures (Fig. 5a). The non-significant degradations over the southern hemisphere at around 950 hPa are likely caused by the terrain height mismatch between the forecasts and the ECMWF analyses. ROMEX outperforms ROMEX20K except for the regions of above 50 hPa over the Southern Hemisphere middle latitudes (Fig. 10e). This is consistent with the detrimental effect of increasing RO data as discussed earlier. Meanwhile, ROMEX and ROMEX20K demonstrate beneficial effects of assimilating additional RO data by showing overall reduced MAER values below 50 hPa, particularly in the tropical tropopause and the middle troposphere at low latitudes. On the other hand, both experiments exacerbate biases in the stratosphere above 50 hPa (Figs. 10b and 10d), which is again attributed to the warming effect introduced by the assimilation of ROMEX observations. Overall, assimilating additional ROMEX RO data improves short-range temperature forecasts in terms of both STDV and MAER when verified against ECMWF analyses, with the exception of the lower stratosphere.



450

451 **Figure 10:** Differences in STDV (a) and (c) and MAER (b) and (d) of the two ROMEX experiments
452 relative to BASE for the 6-h temperature forecasts (unit: K) verified against ECMWF analysis, and
453 difference in STDV (e) and MAER (f) of ROMEX relative to ROMEX20K. Hashed areas overlaid
454 indicate regions of 95% statistical significance level. Positive/negative values of the MAER differences
455 indicate the experiment is farther/closer to the ECMWF analysis.

456

457 Figure 11 presents the STDV and bias of the 6-h specific humidity forecasts of the BASE
458 experiment, verified against the ECMWF analysis (Figs. 11a–b), along with the differences in
459 STDV and MAER between both the ROMEX and ROMEX20K experiments and BASE (Figs.
460 11c–f), and between ROMEX and ROMEX20K (Figs. 11g–h). The STDVs of BASE are
461 largest in the tropical 850 hPa level and decrease gradually with height and toward the poles
462 (Fig. 11a). Both ROMEX experiments show reduced STDV extending from the surface to 300

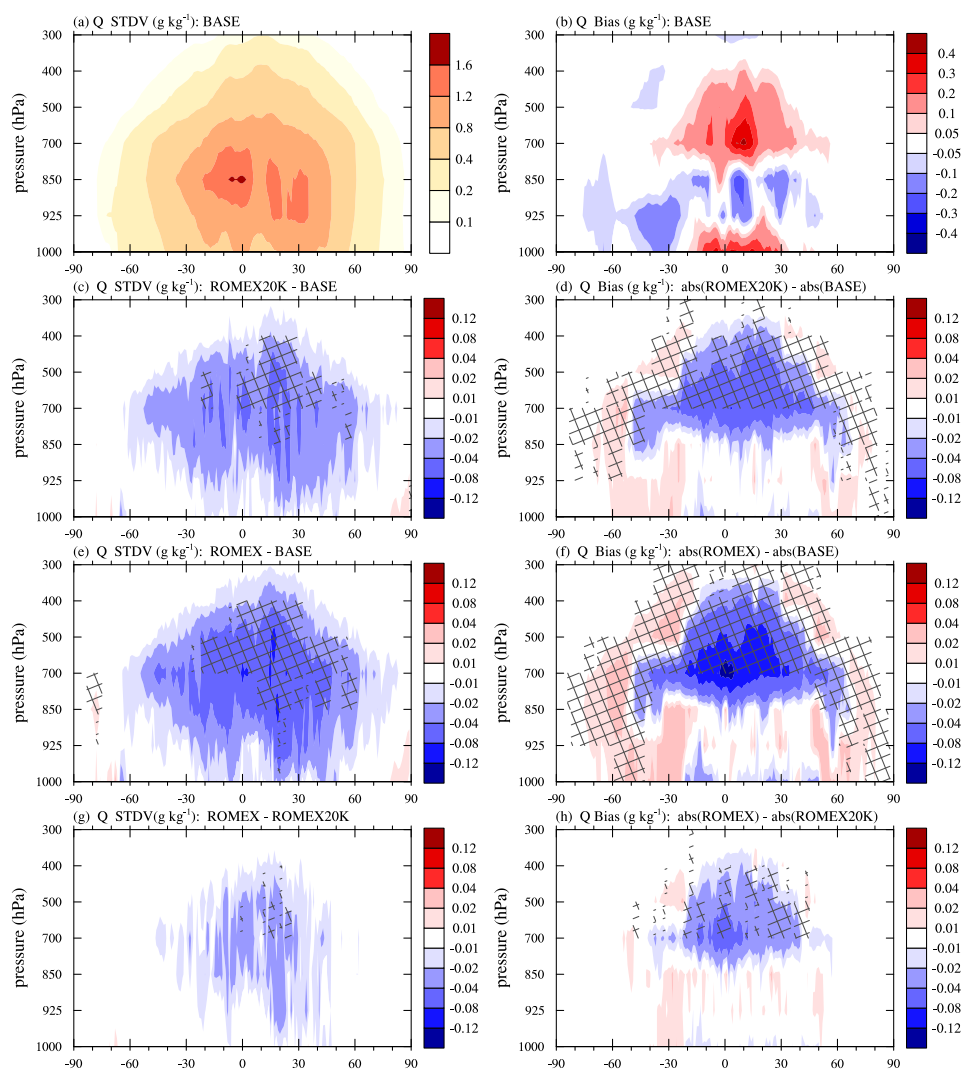


463 hPa (Figs. 11c and e), indicating improved forecast performance. Notably, ROMEX yields
464 additional STDV reductions compared to ROMEX20K in the low-to-mid troposphere over the
465 tropics (Fig. 11g). In terms of bias, BASE exhibits a three-layer structure in the tropics, with
466 positive bias near 700 hPa, negative bias around 900 hPa, and another positive bias near the
467 surface (Fig. 11b). Both ROMEX experiments mitigate the positive bias above 700 hPa (Figs.
468 11d and f), while ROMEX achieves further bias reduction in the 700–400 hPa layer,
469 particularly over tropical regions (Fig. 11h). Overall, both ROMEX and ROMEX20K
470 outperform BASE in terms of humidity forecast skill, with improvements in both STDV and
471 bias, and ROMEX demonstrating an additional advantage over ROMEX20K.

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475

476 **Figure 11:** (a) STDV and (b) bias of BASE 6-h specific humidity forecasts (unit: g kg^{-1}) verified against
 477 ECMWF analysis, and differences in STDV (c) and (e) and MAER (d) and (f) of the two ROMEX
 478 experiments relative to BASE, and difference in STDV (g) and MAER (h) of ROMEX relative to
 479 ROMEX20K. Hashed areas overlaid indicate regions of 95% statistical significance level.
 480 Positive/negative values of such differences indicate the experiment is farther/closer to the ECMWF
 481 analysis.

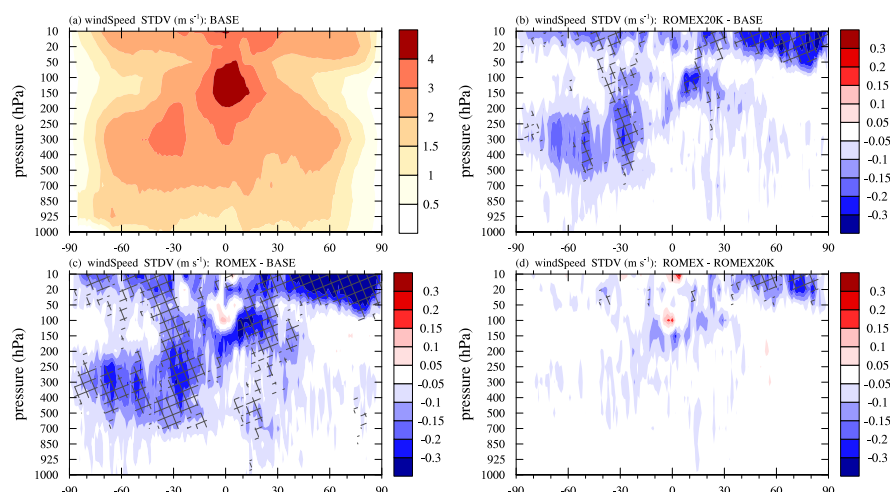


Figure 12: (a) STDV of BASE 6-h wind speed forecasts (unit: m s^{-1}) verified against ECMWF analysis. Differences in STDV for the two ROMEX experiments relative to BASE (b) and (c); and (d) ROMEX relative to ROMEX20K. Hashed areas overlaid indicate regions of 95% statistical significance level. Positive/negative values of such differences indicate the experiment is farther/closer to the ECMWF analysis.

The same diagnostics were also applied to wind speed. Figure 12 displays the 6-h wind speed forecast results: the STDV from BASE (Fig. 12a), the STDV differences between each ROMEX experiment and BASE (Figs. 12b–c), and between the two ROMEX experiments (Fig. 12d). The areas of the largest improvement of wind speed are primarily over the Northern Hemisphere lower stratosphere, the tropical tropopause, and the Southern Hemisphere middle troposphere. Systematic biases are not observed in the wind field and are therefore not presented.

The impact of ROMEX RO assimilation on medium-range forecasts of the JEDI-GFS system is further assessed. Five-day (120-h) forecasts, initiated at each 00Z cycle during September 2022, are examined at 24-h intervals for both experiments. Similar to the short-range evaluation, STDV is calculated against the ECMWF analysis and forecasts of various lead times. The STDV difference with the ECMWF verification is calculated as functions of forecast lead time. Evaluations are conducted in three regions, i.e., the Northern Hemisphere



(NHX, 20°N–80°N), the Tropics (TRO, 20°S–20°N), and the Southern Hemisphere (SHX, 20°S–80°S).

Figures 13–15 illustrate the differences in STDV between ROMEX20K and BASE and between ROMEX and ROMEX20K, as a function of forecast lead time, for temperature, specific humidity, wind speed and geopotential height, respectively. ROMEX20K exhibits reduced STDV, or improved forecast skill, across all three regions throughout the atmosphere above the surface for the temperature 0–5 day forecasts. These reductions persist through the 5-day forecast period, while decaying with lead time (Figs. 13a–c). The most beneficial impacts in the TRO and NHX regions are around 150 hPa, with larger than 8 % improvement at the initial forecast time, and are less than 1 % at day 5 (Figs. 13a–b). In the SHX region, larger than 10 % improvement is observed in a broad layer between 200 and 500 hPa at the initial time, which decays to 2–3% at day 5 (Fig. 13c). With additional data assimilated, ROMEX leads to further improvement in temperature forecast over the lower-to-upper troposphere, lasting up to 3 to 5 days (Figs. 13d–f). The detrimental impacts across the three regions are primarily limited to 20 or 50 hPa with relatively small positive STDV differences toward longer forecast hours for TRO and NHX, and slighter larger values above 50 hPa for SHX than the other two regions. It also shows that the near surface forecast in SHX does not gain benefit from the additional RO assimilation, and the benefit gained above the surface only sustains in the first 3–4 days (Fig. 13f).

ROMEX’s degradation relative to ROMEX20K in the upper levels of SHX extends slightly downward with time. This is closely linked to a warming bias in the lower stratosphere over the mid-latitudes of the Southern Hemisphere, which is introduced by the assimilation of additional RO data (also shown in Figs. 10d, 10f, and 5a).

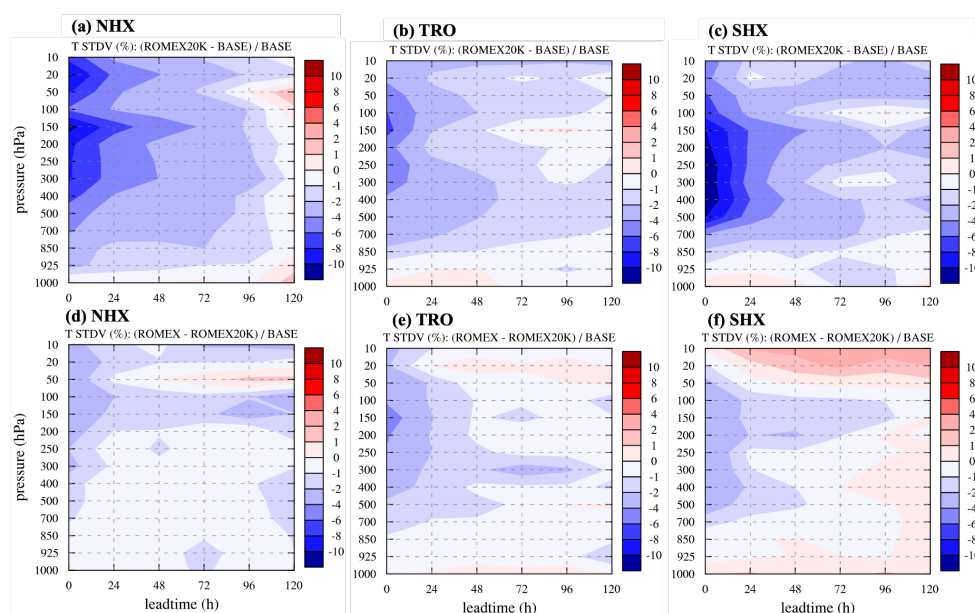


Figure 13: Difference in STDV for temperature forecast of (a-c) ROMEX20K relative to BASE, and (d-f) ROMEX relative to ROMEX20K, verified against ECMWF analysis as a function of forecast lead time for region of (a) and (d) NHX (20°N–80°N), (b) and (e) TRO (20°S–20°N), and (c) and (f) SHX (20°S–80°S).

Similar to the temperature forecasts, the positive impact of assimilating additional ROMEX data on specific humidity forecasts is sustained through five days, with the greatest improvement at the initial time that diminishes rapidly with lead time (Figs. 14a–c). For example, forecast improvements around 500 hPa at the 0-h lead time exceed 6% in all three regions, but decrease to less than 1% at 96-h in NHX and at 72-h in TRO. In SHX, forecast improvements are maintained throughout the troposphere over the 5-day period, with approximately a 2 % reduction in STDV at 120-h. The humidity forecast skill of the increased RO assimilation aligns with that of Prive et al. (2022), in which they stated that the dominant baroclinic process in the SHX winter may account for its longer time scale for improved forecast. As more RO data are assimilated, ROMEX’s positive impact on top of ROMEX20K extends to 5 days in NHX and TRO, but only 2 days in SHX. The relative degradation starts at the surface and propagates upward from day 0 onward (Fig. 14f).

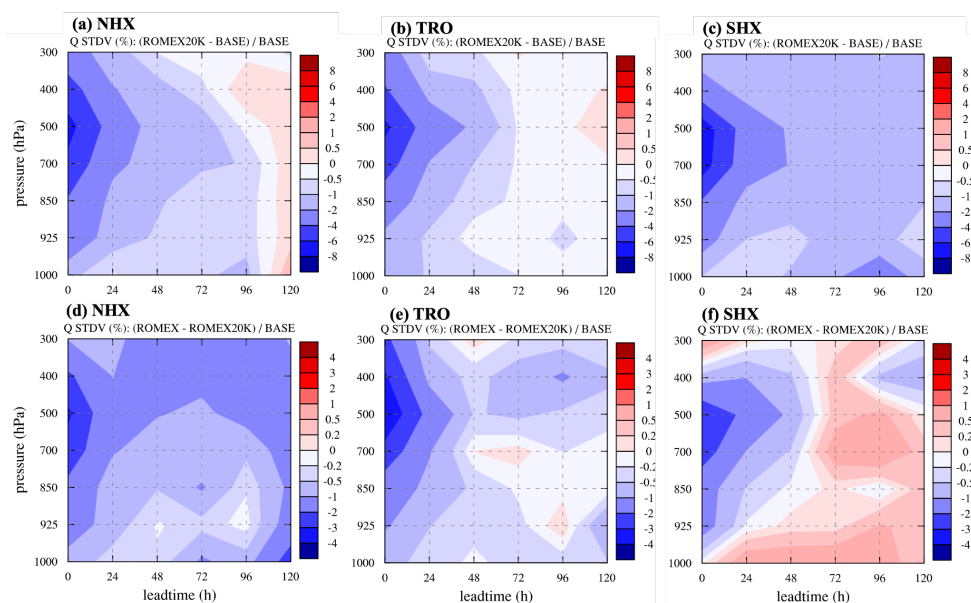


Figure 14: Same as Fig. 13, but for specific humidity forecast.

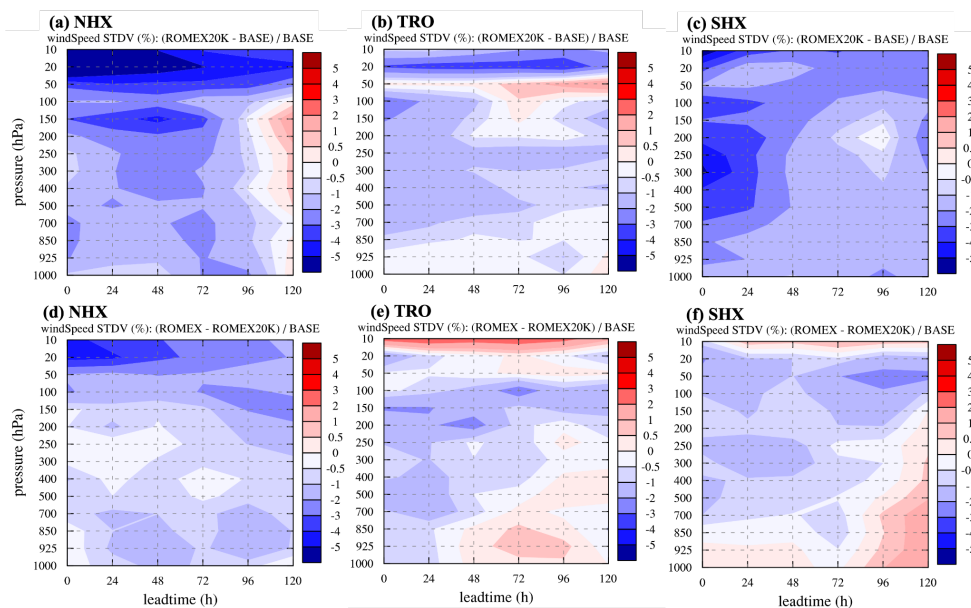


Figure 15: Same as Fig. 13, but for wind speed forecast.



548

549 The impacts of assimilating increased RO data on wind forecast are overall positive from
550 ROMEX20K results for the three regions (Figs. 15a-c). Larger forecast improvements, such as
551 more than 5% STDV reductions in the first three days for NHX, are seen in the lower
552 stratosphere for all regions. Forecast degradations are only noticed after 96-h below 50 hPa
553 (Fig. 15a). TRO are mostly positive as well, except slightly increasing negative impacts around
554 50 hPa beyond 48 h and forward (Fig. 15b). SHX shows all positive impacts from
555 ROMEX20K. Unlike NHX and TRO with the largest positive impacts at upper levels, SHX
556 wind improvements from ROMEX20K are presented for all levels, including 4% improvement
557 between 600–1000hPa for the first 24 h. Also noted is that the impacts of RO data assimilation
558 do not diminish as rapidly as was observed in the temperature and humidity forecasts. For
559 example, the maximum improvement in the NHX troposphere is reached at the 48-h lead time
560 (Fig. 15a).

561 **5. Summary and discussion**

562 This study investigates the impact of increased RO profiles as part of the international
563 collaborative ROMEX project. The current RO profiles available to operational centers are
564 about 8,000–12,000 daily depending on the volume purchased from commercial providers.
565 Earlier studies demonstrated that saturation was not reached with even 128,000 daily profiles.
566 For the first time, the ROMEX project enabled the use of approximately 35,000 RO profiles to
567 explore this further with Observing System Experiments (OSEs).

568 As part of the ROMEX NWP efforts, this study contributes to building consensus on the
569 impact of increased volumes of RO observations and to addressing the risks associated with
570 potential loss of RO capabilities across NWP centers, specifically within the GFS framework
571 – NOAA’s operational forecasting system. At the same time, the study leverages advanced
572 features of JEDI to enhance performance, serving as a valuable platform to evaluate JEDI as
573 the next-generation data assimilation system.

574 Four sets of experiments were conducted over a one-month period in September 2022:
575 noRO, BASE, ROMEX20K, and ROMEX. All experiments assimilated a common set of



576 conventional observations, cloud-motion vectors, and satellite radiances, differing only in the
577 amount of RO data assimilated. The BASE experiment assimilated only the publicly available
578 RO profiles (~8,000 per day), while noRO excluded RO entirely. ROMEX20K and ROMEX
579 assimilated approximately 20,000 and 35,000 daily RO profiles, respectively. The actual
580 number of RO profiles per day varies depending on quality control procedures.

581 The results show that assimilating additional RO profiles significantly improves forecast
582 skill for all key meteorological fields, including temperature, humidity, geopotential height,
583 and wind speed, for most of vertical levels. Forecast improvements were evident in verification
584 against both the critical observations and ECMWF analyses, with impacts lasting up to 5 days
585 (maximum forecast range in the experiments). For example, the STDV reduction of
586 temperature 6h forecasts at 200 hPa, relative to noRO, was 5.3% for ROMEX20K and 6.8%
587 for ROMEX when verified against radiosonde observations. Conversely, withholding RO data
588 led to forecast degradations, with a maximum STDV increase of approximately 0.4 K near 50
589 hPa over the tropics. The results also suggest that forecast improvements scale approximately
590 logarithmically with the number of assimilated profiles, and no evidence of saturation. These
591 results were achieved without any additional tuning of the data assimilation system. All quality
592 control procedures and observation error specifications for RO data used the default, generic
593 configurations implemented in JEDI for testing purposes. The positive outcomes therefore
594 underscore the consistency and robustness of the RO data quality, and demonstrate that
595 assimilating a large volume of RO observations is both feasible and beneficial.

596 However, this effort also revealed areas requiring further investigation. In particular, the
597 assimilation of additional RO increases biases in temperature within the lower stratosphere.
598 The ROMEX20K and ROMEX experiments introduced a cooling effect throughout much of
599 the troposphere and a warming effect above 200 hPa, leading to increased forecast biases
600 relative to the ECMWF analysis. Consequently, Figure 16 illustrates the differences in
601 geopotential height STDV between ROMEX20K and BASE, and between ROMEX and
602 ROMEX20K, as a function of forecast lead time over the Northern Hemisphere. The
603 degradation in height forecasts above 50 hPa due to additional RO data assimilation is clearly
604 evident. Below this level, ROMEX shows forecast improvements lasting up to 5 days (Fig.



16a). The ROMEX versus ROMEX20K comparison (Fig. 16b) suggests that while additional RO data improve forecasts below 50 hPa, they may exacerbate degradation above that level.

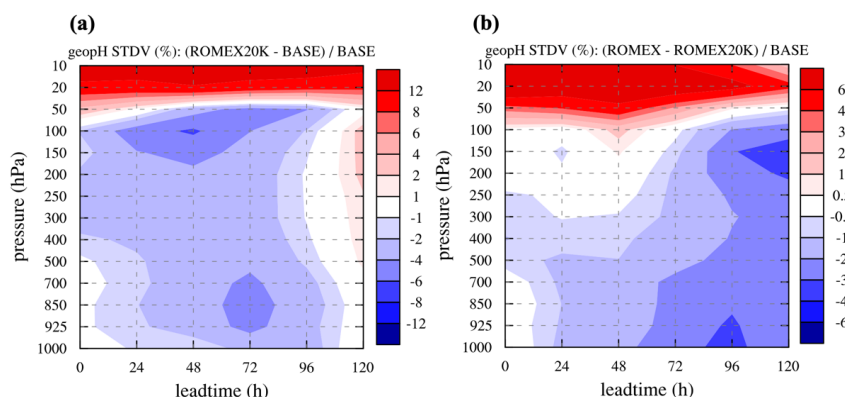


Figure 16: Difference in STDV for geopotential height forecast of (a) ROMEX20K relative to BASE, and (b) ROMEX relative to ROMEX20K, verified against ECMWF analysis as a function of forecast lead time for region of NHX (20°N–80°N).

The sources of these biases are still under investigation. Geopotential height forecast degradation has also been observed by other NWP centers, including the Met Office (Bowler and Lewis 2025) and ECMWF (Lonitz 2025). It is worth noting that the ECMWF analyses used as a reference were produced with the regular volume of RO data assimilated and therefore may not represent the best possible results achievable with the full ROMEX dataset. Additionally, the ECMWF analyses themselves may contain inherent biases, some of which are model-related. Further biases can arise from data processing procedures (e.g. Anthes et al., 2025), and assimilating large volumes of data may amplify such impacts. Moreover, this is the first instance in which the assimilated RO data volume has nearly tripled, and the interactions between RO data and other observations are not yet fully understood.

Ongoing efforts are underway to understand and mitigate these biases. We are actively investigating the role of quality control (QC) and observation errors in the assimilation of ROMEX profiles, especially focusing on whether QC procedures should be regionally adjusted in the upper troposphere and lower stratosphere.



626 To conclude, the assimilation of ROMEX RO data has an overall significantly positive
627 impact in the JEDI-based system. Although no saturation was observed even with the full
628 ROMEX data, the 20K subset significantly improves forecast skill, consistent with
629 recommendations from IROWG-10 (Shao et al. 2025) and the second ROMEX workshop. The
630 combination of the ROPP1D forward operator, the NRL observation error model, and generic
631 quality control, within the JEDI framework, not only enabled the successful assimilation of
632 increased data volume in this study, but also lays the groundwork for future exploration and
633 optimization.

634

635 **Author contribution.** HZ and HS co-designed the experiments, built the end-to-end system,
636 and jointly prepared the manuscript. HZ conducted the experiments and verifications and
637 created the figures. HS and BR prepared the datasets and provided guidance throughout the
638 project. All authors contributed to the interpretation of the results and the review of the
639 manuscript.

640 **Competing interests.** None of the authors has any competing interests.

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649 **Code and Data Availability Statement.** The ROMEX RO profiles are provided through
650 EUMETSAT ROM SAF. Other types of observations are publicly available at the NSF
651 NCAR's Research Data Archive (RDA; <https://rda.ucar.edu/datasets/>), which contains a subset
652 of the NCEP Global Data Assimilation System observations. The JEDI code of the presented
653 work is available at <https://github.com/JCSDA>.



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