

1 **Regionalization of IDF Curves for Mainland China: A Comparative Evaluation of Machine**
2 **Learning versus Spatial Interpolation Techniques**

3 **Yuantian Jiang¹, Wenting Wang^{1*}, Andrew T. Fullhart², Bofu Yu³, Bo Chen⁴**

4 ¹ Department of Geographic Science, Faculty of Arts and Sciences, Beijing Normal University,
5 Zhuhai 519087, China;

6 ² School of Natural Resources and the Environment, University of Arizona, Tucson, AZ, USA;

7 ³ Australian Rivers Institute, School of Engineering and Built Environment, Griffith University,
8 Brisbane, QLD 4111, Australia;

9 ⁴ State Key Laboratory of Earth Surface Processes and Resource Ecology, Faculty of
10 Geographical Science, Beijing Normal University, Beijing 100875, China

11
12 * Corresponding author: Wenting Wang (wangwenting@bnu.edu.cn)

13
14 **Key Points:**

- 15
- 16 • Interpolation and machine learning (ML) were compared for regionalizing Intensity-
Duration-Frequency (IDF) curves for mainland China.
 - 17 • With daily data, ML was able to successfully estimate sub-daily intensities, achieving
18 comparable accuracy using spatial interpolation.
 - 19 • Regionalized IDF curves dataset were developed at 0.1 and 0.5 degrees for mainland
20 China.

21

22 **Abstract**

23 Regionalization of Intensity-Duration-Frequency (IDF) curves is essential for designing
24 stormwater drainage systems, especially in regions without rainfall data of high temporal
25 resolution. However, most studies have not thoroughly compared regionalization methods using
26 sub-daily site observations versus gridded daily precipitation products. The potential of machine
27 learning (ML) methods driven by daily gridded precipitation remains largely underexplored. This
28 study addresses these gaps by regionalizing the IDF curves across mainland China for durations
29 ranging between 1 and 72 hours and return periods ranging from 2 to 1,000 years. Five
30 interpolation methods based on hourly observations from 2363 stations and five machine
31 learning methods using gridded daily dataset were tested for accuracy. Both ML and traditional
32 interpolation methods showed robust performances based on the Kling-Gupta Efficiency (KGE)
33 performance measure. The most successful interpolation method was Kriging with External Drift
34 using mean annual precipitation, with $KGE > 0.96$ for 1-hr-5-yr and 24-hr-5-yr storms and $KGE >$
35 0.84 for 1-hr-100-yr and 24-hr-100-yr storms, while Gradient Boosting was the best-performing
36 ML model, with $KGE > 0.94$ for 1-hr-5-yr and 24-hr-5-yr storms and $KGE > 0.87$ for 1-hr-100-yr
37 and 24-hr-100-yr storms. Notably, even though ML used daily data and interpolation hourly data,
38 the ML accuracy gradually improved, eventually approaching or even surpassing the
39 interpolation methods as the duration and return period increased. Consequently, a regionalized
40 dataset on IDF curves for mainland China with a spatial resolution of 0.1 degrees (and optionally
41 0.5 degrees) was generated using the optimal ML regionalization method.

42 **1 Introduction**

43 The intensity and frequency of extreme precipitation are projected to increase under
44 global warming in response to increased radiative forcing on the hydrologic cycle based on
45 governing physical laws, including the Clausius-Clapeyron relationship (Benestad et al., 2021;
46 Donat et al., 2016; Skliris et al., 2016). These extreme events can lead to disasters, including urban
47 flooding, flash floods, and soil erosion, resulting in significant loss of life and property damage
48 (Gao et al., 2017; Maggioni & Massari, 2019). Intensity-Duration-Frequency (IDF) curves, which
49 define the relationships among rainfall intensity, duration, and frequency of occurrence (or the
50 return period), are derived from historical extreme precipitation events. They play a vital role in
51 the design of stormwater drainage systems and flood protection structures. In regions
52 susceptible to severe flooding, accurately estimating IDF curves is crucial for developing
53 infrastructure that can mitigate property damage and safeguard human lives (Simonović, 2012).

54 Extreme precipitation events are typically characterized by short duration, high intensity,
55 and considerable spatial-temporal variability. As a result, developing accurate IDF curves requires
56 station data with high spatial density and temporal resolution, such as hourly measurements
57 (Förster, 2020; Westra et al., 2014). However, the limited availability of observational stations
58 with high temporal resolution means that IDF curves can only be derived for areas where such
59 stations exist. Consequently, it is challenging to extend these curves to regions lacking high-
60 resolution precipitation records (Sangüesa et al., 2023). With increased urbanization, many new
61 water-related infrastructures will be built in areas without long-term historical precipitation data.
62 Therefore, it is crucial to regionalize and estimate IDF curves for regions where high temporal
63 resolution precipitation data is unavailable or limited.

64 Conventionally, regionalization relies on site-specific observations, with commonly used
65 methods including Regional Frequency Analysis (RFA) and spatial interpolation (Szolgay et al.,
66 2009). Regionalization of station IDF curves through RFA requires identification of homogeneous
67 regions (Schlef et al., 2023). However, this process can be challenging, as defining homogeneous
68 regions often involves subjective judgment and lacks a solid physical basis (Nguyen et al., 2002).
69 As a result, the regions identified may be heterogeneous, which can negatively impact frequency
70 analysis (Halbert et al., 2016; Schlef et al., 2023). Additionally, boundary discontinuities between
71 subregions present another challenge for RFA (Zou et al., 2021).

72 Regionalization methods for IDF curves derived from ground observations also involve
73 interpolation. Two primary approaches are commonly used (Szolgay et al., 2009): (a) directly
74 interpolating the quantiles, moments, or parameters of the distribution function, followed by the
75 calculation of the distribution function for each region; or (b) first calculating the distribution
76 function at the observation stations, then interpolating the return period estimates. Research
77 conducted in the Haihe River Basin, China, for example, shows that interpolating return period
78 estimates are more accurate compared to interpolation based on distribution parameters (Yin et
79 al., 2018).

80 In terms of specific interpolation methods, researchers have frequently employed and
81 compared methods such as Inverse Distance Weighting (IDW), Ordinary Kriging (OK), and Kriging
82 with External Drift (KED). The KED method achieves higher accuracy by incorporating auxiliary
83 variables correlated with the spatial distribution of the target variable (Berndt & Haberlandt,
84 2018; Shehu et al., 2023; Yin et al., 2018). For extreme rainfall prediction, auxiliary variables such
85 as elevation and annual precipitation are correlated with the predicted variable and have the
86 potential to improve interpolation accuracy (Miao et al., 2024; Zou et al., 2021). Notably, the
87 accuracy of interpolation methods fundamentally depends on the quality of the observations. To
88 generate reliable regionalized IDF curves requires a dense observational network with long-term,
89 high-resolution data (Berndt et al., 2014; Papalexioiu, 2018; Papalexioiu & Koutsoyiannis, 2013;
90 Shehu et al., 2023).

91 In recent years, gridded precipitation datasets have undergone significant improvements
92 in both resolution and reliability, enhancing the potential for estimating IDF curves in regions
93 where the deployment of observational stations is challenging (Courty et al., 2019). Consequently,
94 the utilization of gridded precipitation datasets to develop IDF curves has become increasingly
95 prevalent (Ghebreyesus & Sharif, 2021; Mínguez & Herrera, 2023; Noor et al., 2021). These
96 datasets can be generated through various methods, including ground-based station
97 observations, radar, satellite data, or a combination of these sources, resulting in precipitation
98 datasets with varying temporal and spatial resolutions. Gridded IDF products can be estimated
99 using these datasets (Haruna et al., 2024; Lanciotti et al., 2022; Wambura, 2024). However,
100 estimating IDF curves involves assessing precipitation extremes, which may exhibit strong spatial
101 heterogeneity within individual grid cells of precipitation datasets. Further improvements are
102 necessary to enhance the accuracy of gridded IDF products derived from such datasets (Parding
103 et al., 2023; Schilcher et al., 2017).

104 Recent studies have explored the application of machine learning methods for
105 regionalizing IDF curves. Several studies have successfully trained machine learning models on

106 sub-daily precipitation data. For example, one study utilized Support Vector Machine to create a
107 gridded IDF dataset for Canada, while another established a gridded IDF dataset for the Qinghai-
108 Tibet Plateau using Random Forest model (Gaur et al., 2020; Ren et al., 2025). Concurrently, other
109 researches have investigated the non-stationarity of IDF curves under changing climate
110 conditions (Schlef et al., 2023; Vinod & Mahesha, 2024; Zhang et al., 2022). A key challenge in
111 this context is temporal downscaling, where some researches have employed machine learning
112 to disaggregate coarse 3-hour precipitation outputs from climate models into minute-level
113 resolution (Al Kajbaf et al., 2022; Al Kajbaf et al., 2023). This demonstrates the potential of
114 machine learning methods to produce gridded IDF datasets, which traditionally requires sub-daily
115 data, by utilizing precipitation data of a coarser temporal resolution (e.g., daily observations
116 common to rain gauges).

117 The different advantages and disadvantages of the aforementioned IDF curves
118 regionalization methods indicate that it is necessary to identify an optimal regionalization
119 method based on comparative analysis. However, most existing studies assess the accuracy of
120 regionalized IDF curves based on either a site-specific observations or grid-based datasets,
121 without comparing these two kinds of regionalization methods. Additionally, few studies have
122 evaluated the potential of applying machine learning to daily gridded precipitation datasets for
123 estimating regionalized IDF curves, including extrapolation to rainfall intensities at sub-daily time
124 intervals. Furthermore, few studies have computed regionalized IDF curves on a national scale in
125 China.

126 To address these gaps and advance IDF regionalization, this paper is structured as follows.
127 Section 2 describes the study area, data sources, and methods, including the development of IDF
128 curves, spatial interpolation techniques, machine learning approaches, and evaluation metrics.
129 Section 3 presents the results, covering station-level IDF characteristics, a comparative
130 evaluation of regionalization methods at national and regional scales, and the final regionalized
131 IDF curves. Section 4 discusses the implications of the findings, limitation of the final IDF datasets,
132 and opportunities for improvement. Finally, Section 5 summarizes the key conclusions and
133 contributions of the study.

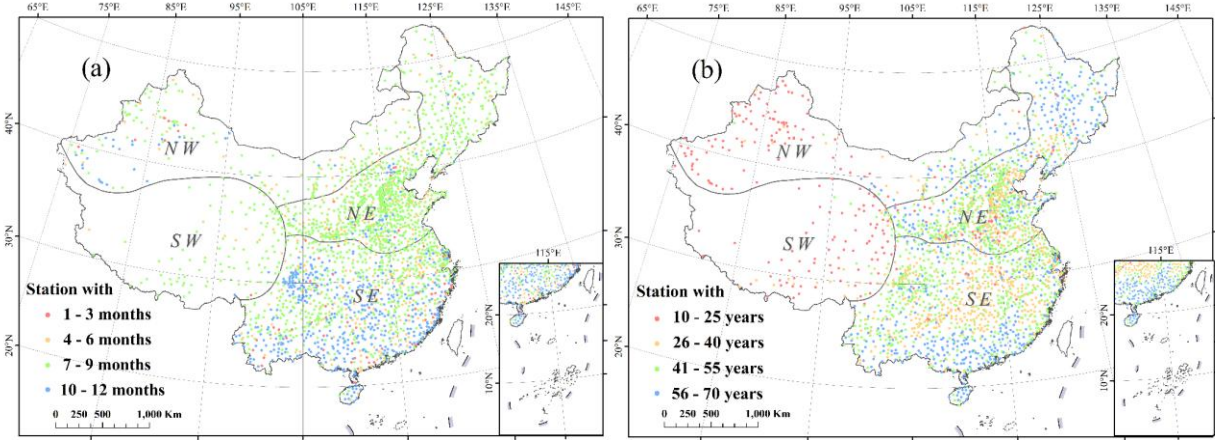
134 **2 Data and Methods**

135 **2.1 Study area and data collection**

136 The study focused on the regionalization of IDF curves on a national scale in mainland
137 China. However, a method that performs well at the national level may exhibit poor accuracy in
138 specific regions. To address this, this study adopted a regionalization scheme that divides
139 mainland China into three distinct regions: the Southwestern Tibetan Plateau (SW) region, the
140 Northwestern Arid (NW) region, and the Eastern Monsoon region (Wang et al., 2019; Zhao, 1983).
141 Due to substantial internal heterogeneity within the Eastern Monsoon region, it was further
142 divided into the Northeastern Monsoon (NE) sub-region and Southeastern Monsoon (SE) sub-
143 region along the Qinling-Huaihe line (Figure 1). This division reflects significant regional
144 differences in climate and topography. To provide additional regional perspective, we also
145 evaluated model performance using an alternative division scheme based on the nine major river
146 basins of China (Figure S3, Tables S23 to S30).

147 To regionalize IDF curves across mainland China, this study utilized two datasets: hourly
148 observations from 2363 stations in mainland China and gridded daily data from the CHM_PRE
149 (Han et al., 2023). The hourly observations were provided and quality-controlled by the National
150 Meteorological Information Center (NMIC) of the China Meteorological Administration (CMA) (Li
151 et al., 2011). Hourly data were observed by siphon self-recording rain gauges, which were not
152 operational during the cold season in some regions to prevent instrument failures caused by low
153 temperatures. The number of operational months for each station is shown in Figure 1a.
154 Additional quality control measures were applied to ensure the reliability of the hourly
155 observations used in developing the IDF curves. Stations were included if they had less than 10%
156 missing data per year. Besides, to ensure a robust sample size for fitting the GEV distribution,
157 only stations with at least 10 years of precipitation records were retained (Ren et al., 2025).
158 Consequently, the data record for the selected stations span from 1951 to 2020, with record
159 lengths ranging from 10 to 70 years, as shown in Figure 1b. To address missing data and prevent
160 underestimation, gaps were handled as follows: (1) for missing periods less than 12 hours, the
161 recorded hourly data were linearly interpolated; (2) for continuous gaps of 12 hours or longer,
162 no interpolation was attempted, and the hours with missing data were assigned a value of zero.
163 From these quality-controlled hourly observations, 27 extreme hourly indices were extracted.
164 Specifically, the annual maximum rainfall for durations from 1 to 24 hours by 1-hour increments,
165 as well as longer durations of 36-hour, 48-hour, and 72-hour, were identified on a yearly basis
166 using a moving window method, resulting in an annual maximum series for each location.

167 The CHM_PRE dataset is a high-quality gridded precipitation dataset covering mainland
168 China. The daily dataset spans the period from 1961 to 2022, with multiple spatial resolutions
169 ($0.1^\circ \times 0.1^\circ$, $0.25^\circ \times 0.25^\circ$, and $0.5^\circ \times 0.5^\circ$). It is constructed using daily observations from 2839
170 gauges across China and nearby regions, employing an interpolation scheme that combines the
171 Parameter-elevation Regression on Independent Slopes Model (PRISM) for the daily climatology
172 field with inverse-distance weighting to interpolate station observations into the ratio field (Han
173 et al., 2023). In this study, $0.1^\circ \times 0.1^\circ$ CHM_PRE was selected. CHM_PRE served as the primary
174 independent variable for the machine learning methods, providing essential input for
175 regionalizing IDF curves across mainland China. Since the dependent variable was based on
176 hourly observations, utilizing independent variables derived from such daily observations did not
177 pose a risk of data leakage. In addition to CHM_PRE, we also explored the CGDPP (Shen et al.,
178 2010) and CHIRPS (Funk et al., 2015) datasets (Tables S1 to S12). However, since they did not
179 demonstrate better performance compared to CHM_PRE, they were not selected for this study.



180
181 **Figure 1.** The spatial distribution of 2363 stations with hourly observations for this study,
182 showing (a) the number of operational months per year, and (b) record lengths. The inset
183 displays the South China Sea Islands within the nine-dash line.

184 2.2 Development of IDF curves

185 The construction of Intensity-Duration-Frequency (IDF) curves in this study follows the
186 Koutsoyiannis method, which calculates generalized rainfall intensity as a function of duration
187 (Koutsoyiannis et al., 1998). Specifically, the generalized intensity i is formulated as:

$$188 \quad i = i_d \cdot b_d \quad (1)$$

189 Here, d denotes the duration in hours, and i_d is the intensity calculated by using the annual
190 maximum series (AMS) method for each duration. The function b_d is:

$$191 \quad b_d = (d + \theta)^\eta \quad (2)$$

192 where the parameters θ and η are estimated for each station by minimizing the Kruskal–Wallis
193 statistic (Koutsoyiannis et al., 1998). Having estimated θ and η , the generalized intensities are
194 pooled together under the assumption that they follow the same distribution for all durations.
195 For the distribution fitting process, this study employs the Generalized Extreme Value (GEV)
196 distribution. The GEV distribution function (Jenkinson, 1955) can be expressed as:

$$197 \quad G(z; \mu, \sigma, \xi) = \exp \left\{ - \left[1 + \xi \left(\frac{z - \mu}{\sigma} \right) \right]^{-1/\xi} \right\}, 1 + \frac{\xi(z - \mu)}{\sigma} > 0, \xi \neq 0 \quad (3)$$

198 where μ is the location parameter, σ is the scale parameter, and ξ is the shape parameter. These
199 parameters can be estimated with the L-moment method (Hosking, 1990). Once the GEV
200 parameters were determined, the corresponding $(1 - r^{-1})$ -th quantile of the GEV distribution
201 for a r -year return period could be calculated. The IDF curves are finally generated by estimating
202 quantiles at specific return periods from the fitted GEV distribution and dividing these quantiles
203 by the term b_d . In this study, the durations selected were whole hours from 1 to 24, along with
204 36, 48, and 72 hours, and the return periods were 2, 5, 10, 20, 50, 100, 200, 500, and 1000 years.

205 2.3 The spatial interpolation methods

206 Traditional spatial interpolation methods were used in this study to regionalize IDF curves.
207 Both deterministic interpolation and geostatistical interpolation methods were utilized. Inverse

208 Distance Weighted (IDW) represents a deterministic interpolation method, while Kriging is a form
 209 of geostatistical interpolation.

210 The fundamental principle of IDW is that the influence of a known point diminishes with
 211 increasing distance from the unknown point. Mathematically, the estimated value $\hat{Z}(x_0)$ at
 212 location x_0 is calculated as follows:

$$213 \quad \hat{Z}(x_0) = \frac{\sum_{i=1}^n \frac{Z(x_i)}{d(x_i, x_0)^p}}{\sum_{i=1}^n \frac{1}{d(x_i, x_0)^p}} \quad (4)$$

214 where $Z(x_i)$ represents the known value at location x_i , $d(x_i, x_0)$ is the Euclidian distance
 215 between the known point x_i and the unknown point x_0 , p is the power parameter that controls
 216 the rate of decay of influence with distance, and n is the number of known points used in the
 217 interpolation. In this study, the IDW method was implemented using a power parameter p of 2,
 218 which is a commonly used value that provides a balance between smoothness and accuracy (Li
 219 & Heap, 2008).

220 Kriging interpolation is a geostatistical method that assumes a stochastic process with
 221 spatial correlation between different points. The core of Kriging interpolation lies in the
 222 calculation of the variogram, which represents the spatial correlation between points at different
 223 distances. Initially, various lag distances are defined, and the semi-variance of residuals at these
 224 lag distances is calculated to generate the empirical variogram. The formula for the semi-variance
 225 (Armstrong, 1998) is as follows:

$$226 \quad \gamma(h) = \frac{1}{2N(h)} \sum_{i=1}^{N(h)} [Z(x_i) - Z(x_i + h)]^2 \quad (5)$$

227 where $\gamma(h)$ is the semivariance at lag distance h , $N(h)$ is the number of pairs of points separated
 228 by h , $Z(x_i)$ is the value at location x_i , and $Z(x_i + h)$ is the value at location $x_i + h$. Subsequently,
 229 a theoretical model is fitted to the empirical variogram. In this study, the commonly used
 230 spherical model was employed.

231 Ordinary Kriging (OK) assumes that the trend is an unknown constant. The formula
 232 (Verworn & Haberlandt, 2011) for Ordinary Kriging is:

$$233 \quad \hat{Z}(x_0) = \sum_{i=1}^n \lambda_i Z(x_i) \quad (6)$$

234 where $\hat{Z}(x_0)$ is the estimated value at location x_0 , λ_i are the weights, and $Z(x_i)$ are the known
 235 values at locations x_i . The weights are determined using the Kriging equations based on the
 236 variogram.

237 Kriging with External Drift (KED) incorporates one or more auxiliary variables as
 238 background information for the interpolation of the primary variable. KED assumes that the
 239 expectation of the primary variable is a linear combination of these auxiliary variables (Verworn
 240 & Haberlandt, 2011):

$$241 \quad E[Z(x) | Y_1(x), Y_2(x), \dots, Y_m(x)] = \beta_0 + \sum_{k=1}^m \beta_k Y_k(x) \quad (7)$$

242 where $E[Z(x) | Y_1(x), Y_2(x), \dots, Y_m(x)]$ is the expectation of the primary variable at location x ,
 243 $\beta_0, \beta_1, \dots, \beta_m$ are the unknown constants, and $Y_1(x), Y_2(x), \dots, Y_m(x)$ are the auxiliary

244 variables. It is important to note that when calculating the variogram using the KED method,
245 $E[Z(x) | Y_1(x), Y_2(x), \dots, Y_m(x)]$ at different spatial points vary. Therefore, the residuals must
246 be calculated first, and then the semi-variance based on residuals is computed. The residuals are
247 calculated as: $R(x) = Z(x) - E[Z(x)]$. In this study, annual precipitation and elevation were
248 used as auxiliary variables (Table 1). The spatial resolution of the annual precipitation data and
249 the elevation data were both $1\text{km} \times 1\text{km}$. The interpolation methods were all performed using
250 the 'gstat' package in R (Pebesma, 2004).

251 2.4 Machine learning methods

252 In this study, machine learning methods were applied using the scikit-learn (sklearn)
253 library in Python (Pedregosa et al., 2011), taking advantage of the CHM_PRE dataset (see Section
254 2.1). For each grid cell, we first calculated a set of independent variables, grouped into four
255 categories: geographic coordinates and elevation, statistics of annual and daily precipitation
256 statistical features, and daily extreme precipitation for different return periods (Table 1). These
257 selected and readily available independent variables without introducing additional
258 meteorological variables allow testing the predictive potential inherent within the precipitation
259 data, while also ensuring the method's practicality and applicability, particularly in regions where
260 other high-quality meteorological data are limited or unavailable. We then aligned each
261 observational station to the grid cell in which the station is located, and treated the station-
262 derived rainfall intensity at specific duration and frequency as dependent target variables. The
263 datasets used to evaluate prediction accuracy were based on all grids that contained both
264 independent and dependent variables.
265

266
267
268

Table 1. Independent variables used in the regionalization methods evaluated. The gridded precipitation data are sourced from CHM_PRE.

Regionalization methods	Category	Description
<i>site-observation-based interpolation method</i>	1. geographic coordinates	latitude longitude
	2. primary variables	Rainfall intensities for 27 durations and 9 return periods based on hourly observations (see Section 2.2)
	3. auxiliary variables	The elevation derived from 1km grid (Figure S1) The mean annual precipitation derived from CHM_PRE (Figure S2)
<i>gridded-precipitation-based machine learning method</i>	1. geographic coordinates and elevation	latitude of the centre of each grid cell longitude of the centre of each grid cell The elevation derived from 1km grid (Figure S1)
	2. annual precipitation statistical features	The mean annual precipitation The standard deviation of annual precipitation The skewness of annual precipitation derived The kurtosis of annual precipitation derived
	3. daily precipitation statistical features	The mean daily precipitation The standard deviation of daily precipitation The skewness of daily precipitation The kurtosis of daily precipitation

The mean annual maximum daily precipitation

The mean annual 95th percentile daily precipitation

The average of the annual sums of daily precipitation that exceeds 95th percentile of daily precipitation for the year

4. daily extreme precipitation

Daily precipitation for 9 return periods (2, 5, 10, 20, 50, 100, 200, 500, and 1000 years). Note that the daily duration is not equivalent to the 24-hour duration, as the latter is based on the moving window method.

269

270 Then a range of four general test cases were extracted from the IDF curves that, for
271 simplicity, are referred to herein as the short duration-small return period (SDSR), short duration-
272 large return period (SDLR), long duration-small return period (LDSR), and long duration-large
273 return period (LDLR) cases, corresponding to the four permutations of IDF cases with 1-hr and
274 24-hr durations, and 5-yr and 100-yr return periods. These targets were selected to represent
275 varying durations and return periods, thereby capturing the different scales of extremes within
276 the IDF framework. These selected cases were used to evaluate the performance of the different
277 regionalization methods.

278 To predict extreme rainfall intensities, 10 different machine learning methods (Table S13)
279 were considered initially, and cross-validation was used to evaluate their accuracy. For each case,
280 we calculated performance scores according to the procedure described in Section 2.5. The five
281 highest-scoring machine learning methods, based on the average scores across all the four cases,
282 were Random Forest (RF), Gradient Boosting (GB), Extremely Randomized Trees (ET), Multilayer
283 Perceptron (MLP), and Linear Regression (LR) (Breiman, 2001; Friedman, 2001; Geurts et al., 2006;
284 Hinton, 1989).

285 It should be noted that we also performed hyperparameter tuning using a grid search for
286 each model for the four cases (Tables S14 to S17). However, the best performance among all
287 machine learning methods with tuned hyperparameters did not show a significant improvement
288 over that with the default settings provided by the Scikit-learn library. Additional experiments
289 using Bayesian optimization (Mockus, 1998; Snoek et al., 2012) to search for optimal
290 hyperparameters also failed to yield notable improvement (Tables S18 to S22). Consequently, for
291 simplicity and consistency, we proceeded with the default hyperparameters throughout the
292 study.

293 **2.5 Evaluation of the regionalization efficiency**

294 This study employed five-fold cross-validation to evaluate the prediction accuracy of
295 interpolation and machine learning methods. Additionally, cross-validation helps mitigate the
296 risk of overfitting, ensuring more robust and generalizable results. In each iteration, 80% of the

297 data was used to train the model, while the remaining 20% was reserved for validation. This
 298 process was repeated five times using sampling without replacement, ensuring that each data
 299 point was used for validation exactly once. The predicted values were then compared with the
 300 observations to evaluate accuracy using the following metrics: Nash-Sutcliffe Efficiency (NSE),
 301 Percent Bias (PBIAS), Root Mean Square Error (RMSE), and Kling-Gupta Efficiency (KGE). The
 302 formulas for these evaluation metrics are provided below:

$$303 \quad NSE = 1 - \frac{\sum_{i=1}^n (O_i - P_i)^2}{\sum_{i=1}^n (O_i - \bar{O})^2} \quad (8)$$

$$304 \quad PBIAS = 100\% \times \frac{\sum_{i=1}^n (P_i - O_i)}{\sum_{i=1}^n O_i} \quad (9)$$

$$305 \quad RMSE = \sqrt{\frac{\sum_{i=1}^n (P_i - O_i)^2}{n}} \quad (10)$$

$$306 \quad KGE = 1 - \sqrt{(CC - 1)^2 + \left(\frac{\sigma_P/\bar{P}}{\sigma_O/\bar{O}} - 1\right)^2 + \left(\frac{\bar{P}}{\bar{O}} - 1\right)^2} \quad (11)$$

307 where O_i denotes the observed values, P_i denotes the predicted values. Specifically, O_i is the
 308 extreme rainfall intensity for a given duration and return period estimated from hourly data at
 309 an individual site, while P_i is the corresponding value for that site predicted by a model (either
 310 interpolation or machine learning) trained on data that excluded the site being evaluated. \bar{O} and
 311 \bar{P} are the mean of the observed and the predicted values, respectively. σ_O and σ_P are the
 312 standard deviations of the observed and predicted values, respectively. CC is the Pearson
 313 correlation coefficient:

$$314 \quad CC = \frac{\sum_{i=1}^n (P_i - \bar{P})(O_i - \bar{O})}{\sqrt{\sum_{i=1}^n (P_i - \bar{P})^2} \sqrt{\sum_{i=1}^n (O_i - \bar{O})^2}} \quad (12)$$

315 The maximum value for NSE is 1, with higher values indicating better predictive
 316 performance. PBIAS values indicate overestimation (positive) or underestimation (negative), with
 317 lower absolute values representing smaller prediction errors. RMSE has a minimum value of 0,
 318 where smaller values denote less error. KGE reflects the correlation, variability, and bias between
 319 predicted and observed values, with its maximum value being 1 and higher values indicating
 320 better performance.

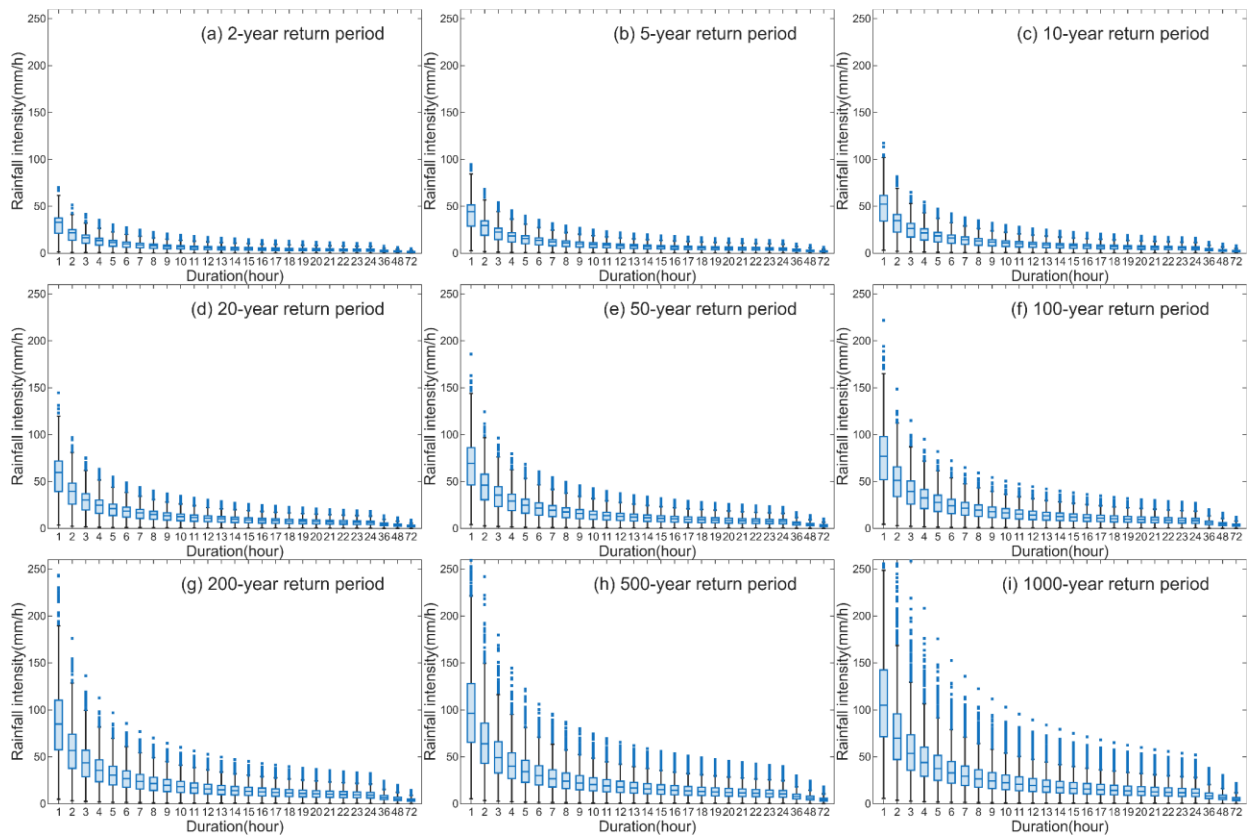
321 To determine the optimal regionalization method for each specific duration and return
 322 period, a weighted composite score based on the above four metrics was defined. First, each
 323 metric was individually standardized using standard deviation normalization among all methods
 324 (mean of 0, with positive or negative values indicating above or below the mean, respectively,
 325 and the values representing multiples of the standard deviation from the mean). The
 326 standardized metrics were then weighted and averaged to calculate the final score for each
 327 method using assigned weights of 0.2, 0.2, 0.2, and 0.4, respectively (RMSE's standardized values
 328 were negated, and PBIAS values were first converted to absolute values before negation). The
 329 method with the highest weighted composite score was selected as the optimal regionalization
 330 method for the given duration and return period. This selection process was performed for the
 331 permutations of 27 durations and 9 return periods ($27 \times 9 = 243$ times). The selected optimal

332 regionalization method was then applied to regionalize extreme rainfall intensity for each specific
333 duration and return period.

334 3 Results

335 3.1 Station-level IDF curves in Mainland China

336 To establish a foundation for regionalizing IDF curves, we first analyzed station-level IDF
337 curves derived from the observed rainfall data across the NMIC ground network. These station-
338 specific IDF curves, serving as the dependent variable in subsequent analyses, provide the most
339 accurate representation of extreme rainfall intensities (Figure 2). The specific statistical
340 characteristics of the four IDF cases used to test regionalization methods are summarized in Table
341 S49. The mean values clearly highlight differences in the magnitude of rainfall intensity across
342 the cases. The LDSR exhibits a mean intensity of 4.5 mm/h, whereas the LDLR shows a
343 significantly higher mean intensity of 8.42 mm/h, indicating higher rainfall intensity. This pattern
344 of increasing rainfall intensity with larger return periods is evident for both short and long
345 durations. In addition, the mean intensity for shorter durations, such as the SDLR (76.38 mm/h),
346 is much higher compared to longer durations like the LDLR. These results emphasize the
347 variations in IDF curves: shorter durations and larger return periods correspond to higher rainfall
348 intensities, underscoring the increasing severity, which can also be inferred to in Section 2.2. Such
349 rainfall intensity of shorter durations and larger return periods may pose additional challenges
350 for accurate prediction.



351

352 **Figure 2.** Boxplots of observed station-level rainfall intensities for various durations and return
 353 periods across stations in mainland China.

354 3.2 Evaluation of regionalization method types

355 To evaluate the performance and reliability of different methods for regionalizing IDF
 356 curves across mainland China, this study systematically compares five traditional interpolation
 357 methods and five machine learning methods. The objective is to assess the predictive capabilities
 358 of these methods, explore the potential of machine learning methods for regionalizing IDF curves,
 359 and demonstrate the reliability of the regionalized IDF curves product created by applying the
 360 optimal method (selected from the ten evaluated methods) to each of the four IDF cases with a
 361 specific duration and return period combination.

362 3.2.1 Interpolation methods

363 Prediction accuracy for the five interpolation methods and five machine learning methods
 364 across four cases are presented in Table 2 and Table 3, respectively. The weighted composite
 365 scores derived from these metrics are shown in Figure 3. Among the interpolation methods,
 366 Inverse Distance Weighting (IDW), a deterministic interpolation technique, consistently
 367 underperformed compared to the geostatistical models represented by the Kriging methods
 368 (Figure 3). This performance gap was particularly evident for cases with smaller return periods.
 369 As shown in Table 2, IDW exhibited inferior performance relative to the Kriging methods across
 370 almost all performance metrics for cases with smaller return periods.

371 Among the four Kriging methods, Kriging with External Drift using annual average
 372 precipitation (KED_AP) demonstrated improved scores across all four cases compared to
 373 Ordinary Kriging (OK) which does not incorporate covariates (Figure 3). However, when elevation
 374 was used as a covariate (KED_DEM), scores declined for all the four cases. Similarly, when both
 375 annual precipitation and elevation were included as covariates (KED_DEM+AP), the pattern
 376 remained consistent with KED_DEM, showing decreased scores for all the four cases compared
 377 to KED_AP (Figure 3). These trends were further corroborated in terms of NSE, PBIAS, RMSE, and
 378 KGE presented in Table 2, which are largely aligned with the observed score patterns.

379 **Table 2.** Accuracy metrics of interpolation methods in mainland China.

	SDSR				SDLR				LDSR				LDLR			
	NSE	PBIAS	RMSE	KGE	NSE	PBIAS	RMSE	KGE	NSE	PBIAS	RMSE	KGE	NSE	PBIAS	RMSE	KGE
IDW	0.93	0.70%	4.22	0.92	0.78	0.71%	15.06	0.82	0.91	0.28%	0.57	0.91	0.81	0.43%	1.69	0.83
OK	0.94	-0.06%	3.92	0.95	0.79	0.06%	14.92	0.82	0.93	-0.01%	0.50	0.94	0.81	0.11%	1.69	0.84
KED_AP	0.94	0.03%	3.79	0.96	0.79	-0.01%	14.72	0.84	0.95	0.07%	0.45	0.96	0.82	0.05%	1.60	0.87
KED_DEM	0.93	0.21%	4.06	0.94	0.78	0.33%	15.24	0.83	0.92	0.07%	0.56	0.93	0.78	0.24%	1.80	0.83
KED_DEM+AP	0.94	0.12%	3.77	0.96	0.79	0.16%	14.89	0.85	0.94	0.12%	0.46	0.96	0.82	0.11%	1.64	0.87

380 *NSE and KGE are dimensionless and RMSE is in mm/h. Positive (Negative) PBIAS (in
 381 percentage) indicates overestimation (underestimation).

382

383 These results indicate that IDW is less effective at predicting extreme rainfall intensities
384 compared to Kriging methods across the four cases. Among the Kriging methods, incorporating
385 annual precipitation as a covariate significantly improved prediction accuracy, whereas including
386 elevation as a covariate did not, and in some cases, even with reduced accuracy. This finding is
387 consistent with previous studies conducted in the Haihe River Basin, China (Zou et al., 2021).

388 3.2.2 Machine learning methods

389 As shown in Figure 3, among the five machine learning methods, Gradient Boosting (GB)
390 performed the best, achieving the highest scores for all cases. In the final regionalized IDF curves,
391 GB also outperformed the other methods and became the optimal method in many cases (Table
392 S31). Following GB’s performance, both Random Forest (RF) and Extremely Randomized Trees
393 (ET) demonstrated relatively stable performance across the four cases, showing acceptable
394 scores with relatively small variability. In contrast, MLP and Linear Regression (LR) exhibited
395 greater performance variability across the four cases (Figure 3). For instance, while MLP achieved
396 high scores for longer durations, its accuracy for shorter durations was notably lower, as
397 indicated by the lower NSE and higher RMSE and PBIAS values (Table 3). However, the strength
398 of MLP’s accuracy in longer-duration cases allowed it to surpass GB and become the optimal
399 method in certain longer-duration cases within the final regionalized IDF curves (Table S31).

400 **Table 3.** Accuracy metrics of machine learning methods in mainland China.

	SDSR				SDLR				LDSR				LDLR			
	NSE	PBIAS	RMSE	KGE	NSE	PBIAS	RMSE	KGE	NSE	PBIAS	RMSE	KGE	NSE	PBIAS	RMSE	KGE
RF	0.92	0.16%	4.34	0.94	0.83	0.37%	13.49	0.87	0.94	-0.03%	0.47	0.96	0.87	0.22%	1.37	0.91
GB	0.92	-0.02%	4.34	0.94	0.83	0.11%	13.29	0.87	0.95	-0.07%	0.44	0.96	0.88	-0.05%	1.36	0.91
ET	0.92	0.10%	4.50	0.93	0.82	0.41%	13.93	0.87	0.93	0.01%	0.50	0.95	0.86	0.18%	1.45	0.90
MLP	0.88	-0.50%	5.36	0.93	0.78	-1.21%	15.10	0.87	0.95	-0.14%	0.45	0.97	0.87	0.18%	1.37	0.91
LR	0.89	0.00%	5.24	0.92	0.80	0.05%	14.56	0.86	0.93	-0.04%	0.50	0.95	0.87	0.00%	1.41	0.91

401 *NSE and KGE are dimensionless and RMSE is in mm/h. Positive (Negative) PBIAS (in
402 percentage) indicates overestimation (underestimation).

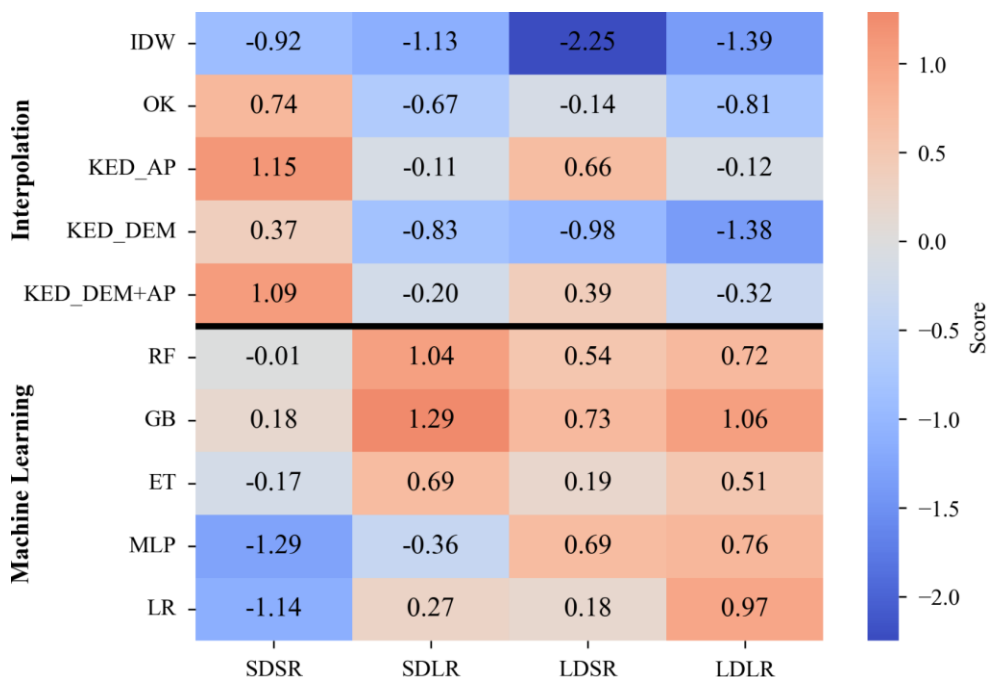
403

404

405 3.2.3 Comparison between interpolation and machine learning methods

406 When comparing interpolation and machine learning methods, the highest composite
407 scores for SDSR, SDLR, LDSR, and LDLR were achieved with KED_AP, GB, GB, and GB, respectively
408 (Figure 3). Notably, the difference in relative scores between the best interpolation and machine
409 learning methods indicate that the performance gap narrows—and may even reverse—as the

410 duration and return period increase (e.g., the score difference is 0.97 for SDSR and -1.18 for LDLR).
 411 This trend indicates that machine learning methods tend to perform better as the duration and
 412 return period increase. This observation is further supported by the final IDF curves dataset
 413 (Table S31). Furthermore, as shown in Table 2 and Table 3, the initial hypothesis in Section 3.1 is
 414 confirmed: shorter duration and larger return periods are usually associated with increased
 415 prediction difficulty and reduced accuracy.



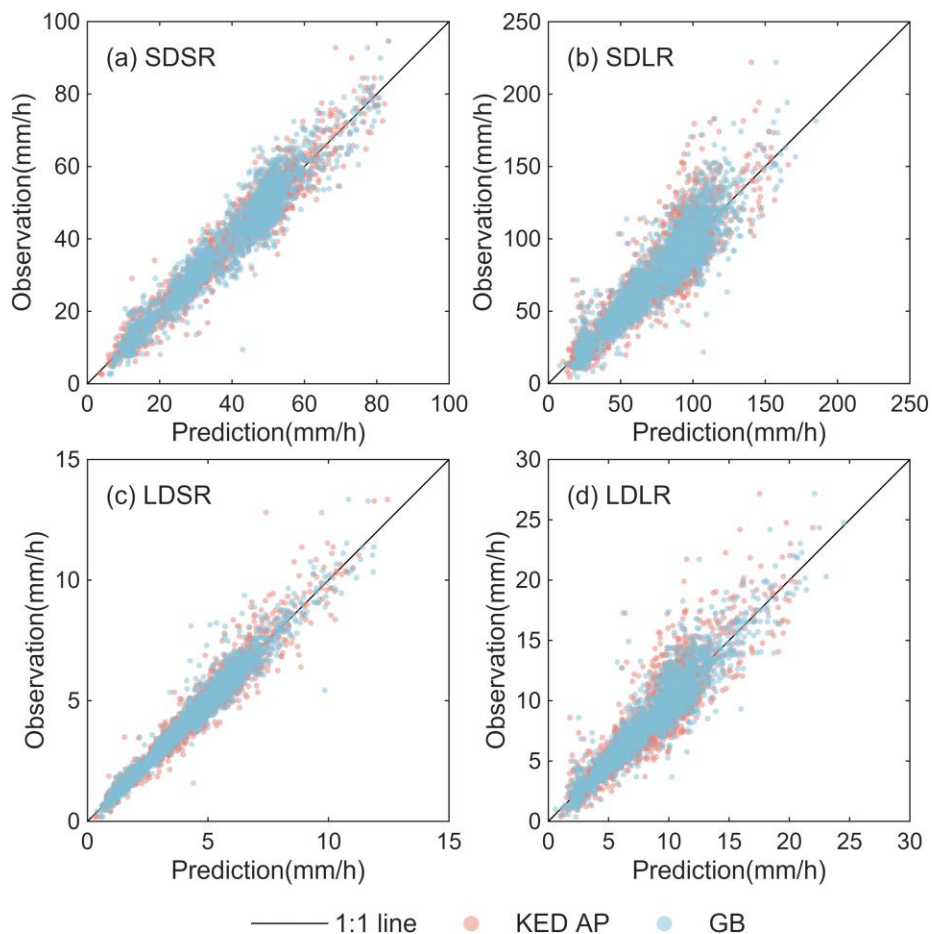
416 **Figure 3.** Weighted composite scores of all regionalization methods. Larger, more positive
 417 values represent better performance, with a value of 0 representing average performance.
 418

419
 420 Methods with the highest scores from both interpolation and machine learning methods
 421 demonstrated strong performance across the four cases. For the most challenging case, SDLR,
 422 the highest scoring methods—KED_AP from the interpolation techniques and GB from the
 423 machine learning models—achieved KGE values not less than 0.84. For SDSR, LDSR, and LDLR, the
 424 highest scoring methods from both interpolation and machine learning methods produced KGE
 425 values of at least 0.94, 0.96, and 0.87, respectively, with PBIAS values close to zero, indicating
 426 little systematic bias (Table 2 and Table 3).

427 In the subsequent sections, this study focuses on KED_AP and GB, which consistently
 428 demonstrated strong and stable performance across the four cases, as representative examples
 429 of interpolation and machine learning methods, respectively. These two methods are used to
 430 show the predictive results at various stations, highlighting both the potential of interpolation
 431 and machine learning methods for IDF regionalization and the regional variations in performance.

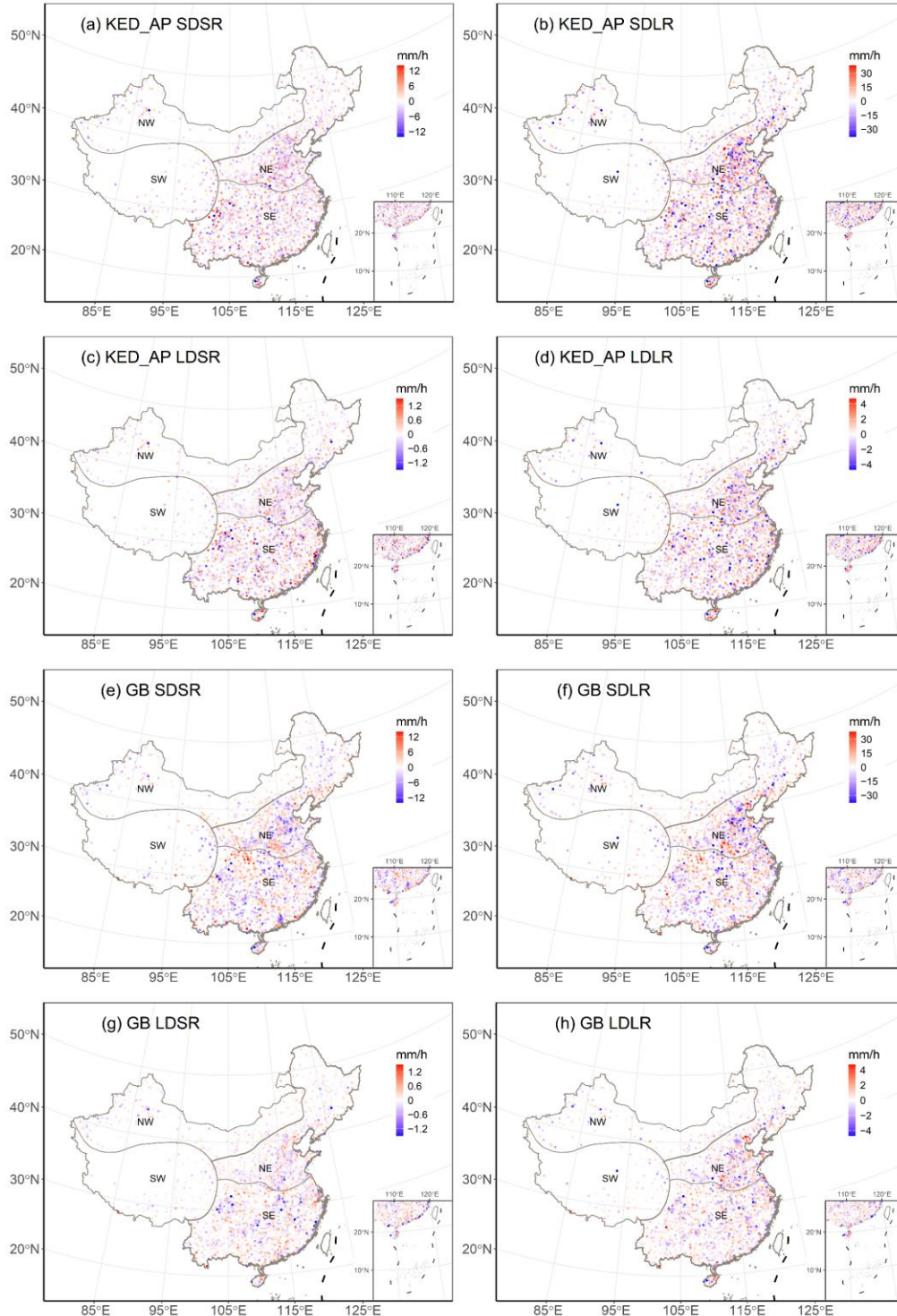
432 Figure 4 presents the prediction results from both interpolation and machine learning.
 433 For all four cases, estimated extreme rainfall intensities are generally aligned along the 1:1 line.
 434 In terms of distribution, the scatter points for the two cases with smaller return periods are more
 435 closely concentrated around the 1:1 line, while those for larger return periods exhibit greater

436 dispersion (Figure 4). In addition, analysis of relative errors (Figure S5) reveals comparable
 437 performance between the two methods across all cases. The relative errors are notably smaller
 438 for smaller return periods (median RE: 0.39-0.65%) compared to larger return periods (median
 439 RE: 1.5-2.5%), indicating increased prediction uncertainty for more extreme events.



440
 441 **Figure 4.** Scatterplots of observations and predictions for KED_AP (the best interpolation
 442 method) and GB (the best machine learning method).

443 The residuals (calculated as the difference between predicted and observed values) for
 444 both methods mostly show distributions that are close to be normally distributed (Figure S6).
 445 However, compared to the true normal distribution, the residual distributions still exhibit mild
 446 negative skewness and higher kurtosis, suggesting that both methods tend to underestimate
 447 certain extreme values. These underestimation is evident in Figure 4, where the corresponding
 448 values are scattered above the 1:1 line. For an ideal model, the residuals should be randomly
 449 distributed across the study area, with no obvious spatial patterns influenced by location or the
 450 magnitude of observed values. The spatial distribution of residuals is illustrated in Figure 5. The
 451 overall spatial distribution of residuals is relatively random, with no obvious clustering or trend.
 452 Consistent with the findings from Figure S6, there are marginally more instances of severe
 453 underestimation than severe overestimation across all cases. The above results of residual
 454 analysis also indicate that the predictive performance of interpolation and machine learning is
 455 satisfactory, with both providing reliable predictions across mainland China.



456
 457 **Figure 5.** Spatial distribution of the residuals for KED_AP (the best interpolation method, (a)-(d))
 458 and GB (the best machine learning method, (e)-(h)). Blue points represent underestimation,
 459 while red points represent overestimation.

460 In short, both interpolation and machine learning methods demonstrate comparable and
 461 robust performance across various cases, whether for short or long durations and return periods.

462 This consistent accuracy in terms of accuracy metrics, residuals, and relative errors highlights the
463 strong potential of the two methods for the regionalization of IDF curves in mainland China.
464 Notably, machine learning, as a method not previously applied to IDF curves regionalization in
465 mainland China, demonstrates accuracy comparable to traditional interpolation methods and
466 even outperforms them in predicting cases for long durations and large return periods.
467 Additionally, machine learning offers distinct advantages, such as requiring only daily
468 precipitation data for regionalization once a model is trained (in contrast with interpolation
469 methods, which always require known sub-daily level target values). Furthermore, machine
470 learning is independent of station setup conditions, and possesses strong adaptability for future
471 updates, forecasting, and climate projection. These features suggest that machine learning has
472 the potential to replace interpolation methods in future IDF curves regionalization applications.

473 3.3 Regional performance evaluation of interpolation and machine learning methods

474 A method that performs well at the national level may exhibit poor accuracy for specific
475 regions. To address this, this study adopted a regionalization scheme that divides mainland China
476 into four regions (See Section 2.1). Using this scheme, we conducted a regional assessment across
477 mainland China, using KED_AP and GB as representatives of interpolation and machine learning
478 methods, respectively, to explore their applicability in the four distinct regions. Notably, we did
479 not retrain the models for regional scales but used the nationally trained models to compute
480 accuracy metrics within each region.

481 As an interpolation method, the predictive accuracy of KED_AP is influenced by station
482 density and whether the interpolation assumptions, such as decreasing correlation with distance,
483 are satisfied. In the NE region, where station density is high and precipitation patterns are
484 relatively stable, KED_AP achieves high predictive accuracy, with a KGE of 0.94 for SDSR and LDSR
485 (Table 4). In contrast, in the SE region, characterized by a humid subtropical and monsoon climate,
486 the performance of KED_AP declines for SDSR, SDLR, and LDSR compared to the NE region,
487 despite the dense station network (Table 4). This decline is likely due to the more extreme nature
488 of precipitation in the SE region, which may gradually violate the assumption of decreasing
489 correlation with distance. In the NW region, precipitation extremes are less pronounced due to
490 its inland location, but the low station density reduces predictive accuracy compared to the NE
491 region (Table 4). Lastly, in the SW region, KED_AP performs significantly worse across all four IDF
492 cases, with the highest KGE for LDSR reaching only 0.72 (Table 4). As shown in Figure S1, the
493 poor accuracy in the SW region may be attributed to the sparse station density and the region's
494 complex topography, which greatly affects interpolation accuracy, consistent with findings in
495 previous studies (Miao et al., 2024).

496 The machine learning method, GB, demonstrates notable adaptability across most
497 regions and is often comparable to or better than interpolation methods, particularly for longer
498 durations and return periods. As shown in Table 4, in the NE region, GB achieves a KGE of 0.92
499 for SDSR, closely matching KED_AP's performance of 0.94. For LDSR, GB's KGE reaches 0.95,
500 slightly exceeding that of KED_AP. In the SE region, although GB slightly underperforms KED_AP
501 for shorter durations and return periods (e.g., KGE of 0.86 for SDSR compared to 0.88 for KED_AP),
502 its relative performance improves as duration and return period increase. For example, for LDLR,
503 GB achieves a KGE of 0.87, surpassing KED_AP's KGE of 0.79. Similarly, in the NW region, GB's

504 performance for LDLR reaches a KGE of 0.77, exceeding KED_AP's KGE of 0.68. However, in the
 505 SW region, GB's performance is considerably lower than that in other regions across all four IDF
 506 cases, with the highest KGE reaching only 0.65. This underperformance is likely attributable to
 507 the pronounced terrain undulation and sparse station distribution in the SW region, where the
 508 complex precipitation patterns are challenging for a nationally trained GB model to capture
 509 accurately. Additionally, the representativeness errors in the gridded precipitation data may
 510 further hinder the accuracy of the machine learning method in this region. Overall, the results
 511 demonstrate that GB exhibits notable adaptability across most regions, with its predictive
 512 accuracy relative to interpolation methods improving as duration and return periods increase,
 513 thereby demonstrating the robustness of this trend at the national level.

514

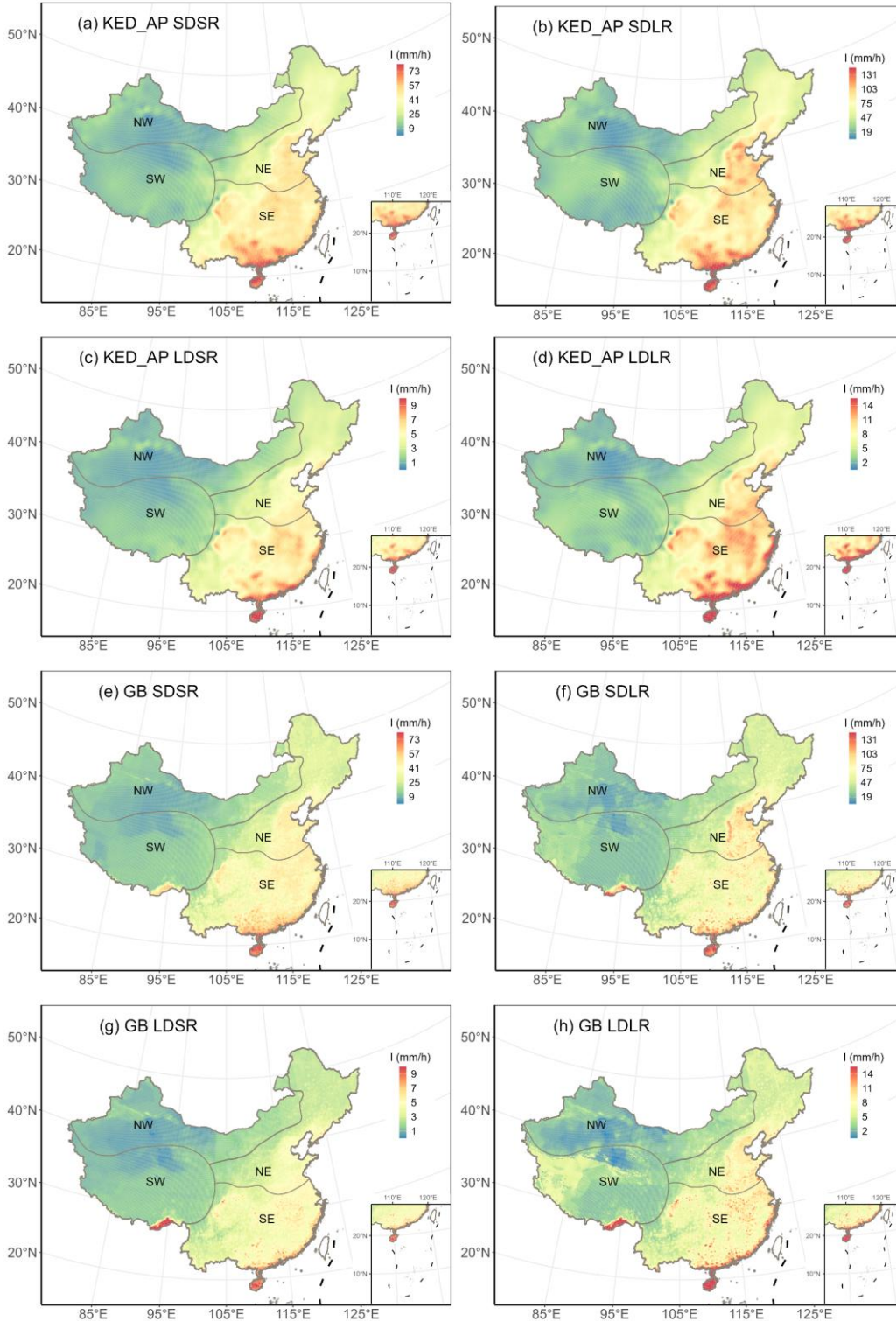
515 **Table 4.** Regional accuracy metrics for KED_AP (the best interpolation method) and GB
 516 (the best machine learning method).

	SDSR				SDLR				LDSR				LDLR			
	<i>NSE</i>	<i>PBIAS</i>	<i>RMSE</i>	<i>KGE</i>	<i>NSE</i>	<i>PBIAS</i>	<i>RMSE</i>	<i>KGE</i>	<i>NSE</i>	<i>PBIAS</i>	<i>RMSE</i>	<i>KGE</i>	<i>NSE</i>	<i>PBIAS</i>	<i>RMSE</i>	<i>KGE</i>
<i>The interpolation method: Kriging methods incorporating annual precipitation as a covariate (KED_AP)</i>																
NE	0.92	-0.10	3.32	0.94	0.73	-0.06	15.12	0.78	0.92	-0.11	0.32	0.94	0.71	-0.32	1.54	0.77
NW	0.79	-0.50	2.98	0.87	0.47	-0.74	10.46	0.63	0.86	0.07	0.26	0.90	0.55	-0.79	0.98	0.68
SW	0.46	0.06	3.05	0.65	0.13	0.55	9.47	0.31	0.52	0.81	0.25	0.72	-0.12	2.84	1.10	0.03
SE	0.85	0.14	4.30	0.88	0.65	0.05	15.46	0.72	0.88	0.15	0.56	0.91	0.74	0.26	1.78	0.79
Mainland China	0.94	0.03	3.79	0.96	0.79	-0.01	14.72	0.84	0.95	0.07	0.45	0.96	0.82	0.05	1.60	0.87
<i>The machine learning method: Gradient Boosting (GB)</i>																
NE	0.88	-0.12	3.92	0.92	0.77	0.06	13.94	0.83	0.93	-0.14	0.32	0.95	0.78	-0.09	1.36	0.84
NW	0.83	1.40	2.73	0.90	0.57	1.82	9.40	0.69	0.89	1.16	0.23	0.94	0.66	2.16	0.86	0.77
SW	-0.27	3.57	4.69	0.29	-0.59	5.92	12.86	0.14	0.17	4.33	0.33	0.65	-0.18	4.11	1.13	0.35
SE	0.82	-0.11	4.83	0.86	0.74	-0.13	13.41	0.8	0.89	-0.23	0.55	0.92	0.83	-0.26	1.46	0.87
Mainland China	0.92	-0.01	4.34	0.94	0.83	0.10	13.30	0.87	0.95	-0.08	0.44	0.96	0.88	-0.05	1.36	0.91

517

518 3.4 The regionalized IDF curves in Mainland China

519 Figure 6 shows the spatial distribution of four cases predicted by KED_AP and GB. Both
520 methods reveal a spatial pattern with a decreasing trend from the southeastern coastal areas to
521 the northwestern interior, with overall similar distributions. This distribution is consistent across
522 different durations and return periods, highlighting the influence of China's varied climatic and
523 topographic conditions. In the SE region, which is characterized by humid subtropical and
524 monsoon climates, higher extreme rainfall intensities are evident. This region experiences the
525 most intense rainfall due to its proximity to oceanic moisture sources, frequent passage of
526 typhoons, and summer monsoon systems. The reduction in extreme rainfall intensity moving
527 inland reflects the weakening of moisture-laden winds as they traverse the country and
528 encounter various geographical barriers, such as the Tibetan Plateau. This trend aligns with
529 another study on design rainstorms with a 2-year return period and a 20-minute duration in China,
530 which shows significantly higher rainstorm amounts in the east compared to the west, and in the
531 south compared to the north (Shao & Liu, 2018). However, notable local differences between
532 both methods can be observed. The interpolated rainfall intensity appears smoother and more
533 uniform. In contrast, the machine learning method exhibits weaker local autocorrelation,
534 resulting in more abrupt spatial changes(Figure 6). Interpolation relies on distant stations, leading
535 to values that are closer to those observed at existing stations, producing a more homogeneous
536 spatial distribution. The machine learning method, however, is more influenced by local
537 geographical conditions, resulting in greater spatial heterogeneity and highlighting the diversity
538 of the terrain and climate.



539
540
541
542
543

Figure 6. Spatial distribution of predicted rainfall intensity (mm/h) from KED_AP (the best interpolation method, (a)-(d)) and GB (the best machine learning method, (e)-(h)).

Notably, the machine learning method predicts local peaks in the southeastern edge of the SW region, which are not captured by the interpolation results (Figure 6). This discrepancy

544 was also found in comparisons of rainfall erosivity (i.e., the potential for water erosion caused by
545 rainfall) estimated with Global Climate Models (GCM) generated rainfall and gauge-observed
546 hourly rainfall (Wang et al., 2023), which likely arises from the influence of orographic rainfall on
547 the windward slopes, leading to more extreme precipitation in this area. The machine learning
548 method can detect this feature from the grid data. In contrast, the interpolation method,
549 constrained by the sparse distribution of rain gauges, extrapolates the rainfall intensity from
550 surrounding areas with less rainfall, thereby missing this localized phenomenon. This underscores
551 the advantage of machine learning methods based on gridded precipitation data in regions where
552 station data is sparse or unavailable.

553 **4 Discussion**

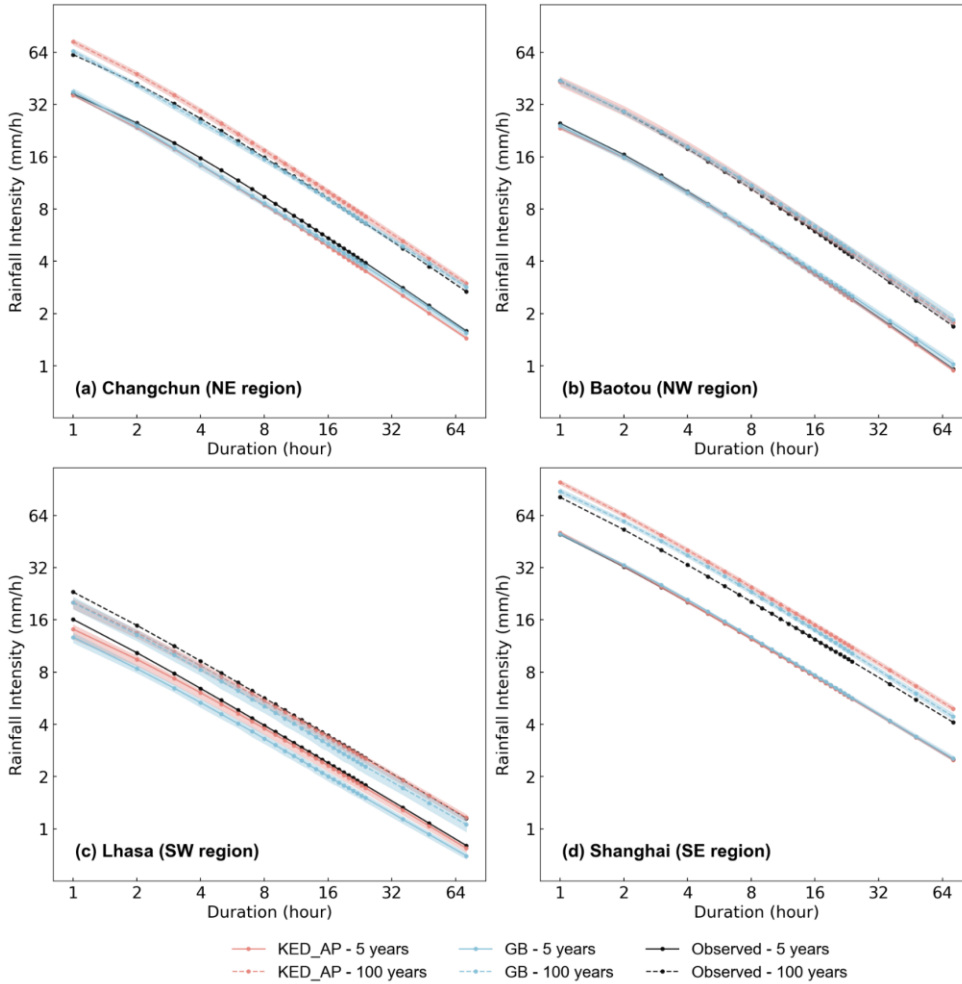
554 Effectiveness of different regionalization methods was comprehensively evaluated in this
555 study for estimating regionalized IDF curves across mainland China. Specifically, the IDF curves
556 were regionalized using five traditional interpolation methods—Inverse Distance Weighted
557 (IDW), Ordinary Kriging (OK), Kriging with External Drift using annual average precipitation
558 (KED_AP), Kriging with External Drift using elevation (KED_DEM), and Kriging with External Drift
559 using both elevation and annual average precipitation (KED_DEM+AP)—along with five different
560 machine learning methods: Random Forest (RF), Gradient Boosting (GB), Extremely Randomized
561 Trees (ET), Multilayer Perceptron (MLP), and Linear Regression (LR). To evaluate the accuracy,
562 each regionalization method was compared against site-specific IDF curves used as benchmarks.
563 In addition, this study attempted to test whether any of the methods combined daily
564 precipitation data could be satisfactorily used to regionalize and extrapolate IDF curves for storm
565 durations much shorter than 24 hours.

566 The comparison of interpolation and machine learning methods for IDF curves across
567 mainland China reveals several notable patterns and insights. Our analysis demonstrates that
568 both methods can effectively estimate extreme rainfall intensities, with the best-performing
569 methods from each category—KED_AP for interpolation and GB for machine learning—achieving
570 comparable accuracy levels. Furthermore, the relationship between prediction accuracy and IDF
571 cases exhibits clear patterns: both methods show decreased performance with shorter durations
572 and larger return periods. This pattern aligns with the inherent challenges in predicting more
573 extreme and temporally concentrated rainfall events. KED_AP outperforms other interpolation
574 methods, such as OK and KED_DEM, likely because the annual average precipitation is a more
575 effective covariate for capturing the spatial patterns of extreme rainfall across diverse terrains
576 compared to other covariates considered in this study. This advantage may stem from its strong
577 correlation with extreme precipitation, whereas elevation, despite some correlation with
578 extreme precipitation in regions of complex topography, appears less effective in flatter regions
579 where its influence on extreme precipitation diminishes (Zou et al., 2021). As for machine
580 learning methods, ensemble learning techniques, particularly GB, have demonstrated stable and
581 excellent performance. To assess the interpretability of the GB model, a feature importance
582 analysis using Shapley Additive Explanations (Lundberg & Lee, 2017) was conducted (Figure S4).
583 This analysis reveals that the model's predictive power is primarily driven by features derived
584 from the gridded precipitation data characterizing daily extremes, specifically the gridded daily
585 precipitation of varying return periods and the average annual maximum daily precipitation. This

586 is likely because the model effectively learns the strong correlation between extreme
587 precipitation events at sub-daily scales and corresponding daily gridded precipitation extremes,
588 capturing this shared pattern across numerous samples through its non-linear capabilities.
589 Furthermore, geographic variables like latitude, longitude, and elevation consistently rank as
590 important secondary features. These variables establish the background precipitation gradient
591 controlled by climatic and topographic factors across mainland China. This spatial baseline,
592 together with daily extreme precipitation features, enables the model to produce spatially
593 heterogeneous IDF estimations.

594 The national-level analysis, however, raises an important question about the spatial
595 consistency of these findings. While both methods demonstrate robust overall performance,
596 their effectiveness may vary across the diverse geographical and climatic regions in mainland
597 China. This consideration is particularly relevant given the country's complex topography and
598 rainfall patterns. Therefore, it is essential to examine how these methods perform across
599 different regions. The analysis in Section 3.3 reveals that both methods achieve acceptable
600 accuracy in the NE, SE, and NW regions, while in the SW region, both methods show reduced
601 performance, with the machine learning method exhibiting lower accuracy than the interpolation
602 method (Table 4). Figure 7 further illustrates IDF curves from representative stations in in each
603 of the four regions, indicating that both the interpolation and machine learning methods produce
604 IDF curves with low uncertainty and high consistency with observed curves. This further
605 corroborates the reliability of both methods. The performance discrepancy between the
606 interpolation and machine learning methods in the SW region is likely because, despite the
607 overall lower station density and complex terrain in the SW region, the training and validation
608 stations are relatively concentrated, thereby reducing the limitations imposed by station density
609 on interpolation accuracy (Figure 1). This suggests that the interpolation method may exhibit
610 reduced predictive accuracy in the western areas of the SW region, where validation stations are
611 sparsely distributed, although this conclusion cannot be confirmed due to the absence of
612 observational data for verification. In comparison, the accuracy of the machine learning method
613 is more constrained by the availability of local samples. The machine learning model, trained on
614 national station data, cannot accurately capture the unique rainfall patterns in the SW region due
615 to the limited number of local training stations there. In addition, the representativeness errors
616 inherent in gridded precipitation data could further impair the machine learning method's
617 predictive accuracy in this region.

618



619

620 **Figure 7** IDF curves using KED_AP (interpolation method), GB (machine learning method),
 621 and observed data for (a) Changchun in the NE region, (b) Baotou in the NW region, (c) Lhasa in
 622 the SW region, and (d) Shanghai in the SE region, with return periods of 5 and 100 years. For all
 623 stations in the mainland China except the four shown, 80% were randomly selected as training
 624 samples for both interpolation and machine learning methods, with this process repeated 500
 625 times using Monte Carlo sampling to predict the IDF curves for these four stations. Black lines
 626 represent the observed IDF curves, red lines represent the mean of the 500 samples for the
 627 interpolation method, and blue lines represent the mean for the machine learning method, with
 628 solid lines indicating the 5-year return period and dashed lines indicating the 100-year return
 629 period. Shaded areas correspond to the 2.5th and 97.5th percentiles of the 500 samples.

630

631 A particularly interesting finding is the gradual shift with duration and return period in
 632 relative performance between the two methods evaluated. As duration and return period
 633 increase, the accuracy of machine learning in most regions gradually improves, eventually
 634 approaching or even surpassing the interpolation, consistent with the national evaluation (Table
 635 4 and Figure 3). The enhanced relative performance of machine learning methods for longer
 636 durations may be attributed to the disparity in the temporal resolution of the independent

637 variables (Table 1). Specifically, the interpolation utilizes station rainfall data with higher
638 temporal resolution (hourly), whereas the machine learning relies on daily gridded precipitation
639 data. Consequently, the accuracy of the interpolation remains relatively unaffected by shorter
640 durations, whereas machine learning exhibits reduced predictive capability for short-duration
641 rainfall events but demonstrates its inherent predictive advantages for long-duration rainfall.
642 This suggests that utilizing gridded precipitation data with higher temporal resolution as
643 independent variables would likely lead to improved accuracy in regionalized IDF curves based
644 on machine learning. Furthermore, the enhanced relative performance of machine learning
645 methods for larger return periods may stem from their ability to better handle extreme rainfall
646 events, where the assumptions of spatial autocorrelation required by interpolation methods
647 become less tenable. Machine learning models make better use of local precipitation and
648 topographical features, allowing them to capture the complex non-linear relationships between
649 these features and extreme rainfall more effectively than interpolation methods.

650 Further practical differences between interpolation and machine learning methods were
651 identified that generally apply in various contexts. The machine learning approach is useful in
652 applications where a trained model may be used to make predictions for areas with observations.
653 This means that machine learning is particularly useful for forecasting of future prediction
654 surfaces, while in contrast, forecasting is not possible with interpolation techniques that always
655 require point observations. By comparison, interpolation techniques may be preferable for
656 producing prediction surfaces for historical time frames where observations are available. The
657 geospatial approach used by most interpolation methods considers data in the surrounding
658 neighborhood of the point for prediction. Consequently, interpolation methods may be expected
659 to result in smoother surfaces with a higher degree of spatial autocorrelation compared to the
660 test machine learning methods that only consider information where prediction is attempted.
661 Therefore, the choice of prediction method may depend on the context of the application and
662 the available data.

663 Several limitations regarding data availability and methodological assumptions should be
664 noted. While our study found that the accuracies of machine learning and interpolation
665 accuracies are comparable, this conclusion may change as more validation stations are included.
666 For instance, given that interpolation accuracy is constrained by station density, large areas in
667 the NW and SW regions both lack stations and are distant from existing stations (Figure 1), likely
668 resulting in severely affected interpolation accuracy in these areas. Similarly, since the accuracy
669 of machine learning is constrained by the availability of non-local samples, and the gridded
670 precipitation datasets in these regions may have been significantly underestimated (Miao et al.,
671 2024), the results derived from machine learning based on these datasets may also be severely
672 affected. Due to the lack of stations for verification, these remain unconfirmed and will need to
673 be validated with installation of additional stations in the future. It is also important to note that,
674 although regionalized IDF were generated for the NW and SW regions in this study, their use
675 should be approached with caution due to the lack of station-based verification. Users are
676 strongly advised to consider the higher uncertainty associated with predictions in these areas
677 and validate results against local observations if available. For the SW region, an alternative high-
678 resolution IDF dataset ($1/30^\circ$) trained with local samples offers a valuable supplement to our
679 work (Ren et al., 2025). While our study focuses on the robustness and general applicability of

680 regionalization methods at a near-national scale across diverse climates and topographies, their
681 study concentrates on achieving high-resolution regionalization within one of the most complex
682 regions in mainland China. Another limitation worth noting is that our study employed stationary
683 IDF curves derived from historical observations, assuming temporal invariance in the extreme
684 precipitation distribution. With climate change both the frequency and intensity of extreme
685 precipitation events are expected to increase, thereby affecting the distributional parameters,
686 future studies should investigate how these regionalization methods perform under non-
687 stationary conditions. Furthermore, limitations inherent in the observational records must be
688 acknowledged, specifically the uncertainty associated with extrapolating 10-70 years of records
689 to 1000-year return periods, and the exclusion of sub-hourly events due to the minimum input
690 duration of one hour.

691 Finally, this study not only highlights the promise of machine learning but also points to
692 opportunities for further refinement. The potential for improving the accuracy of interpolation
693 methods in this study area is limited, due to the lack of higher station density and finer temporal
694 resolution observations. Conversely, machine learning methods have substantial room for
695 improvement through enhanced temporal and spatial resolution of gridded precipitation data,
696 improved data quality, incorporation of additional meteorological variables such as temperature,
697 relative humidity, and atmospheric circulation patterns, as well as the development of more
698 advanced machine learning techniques and exploration of more comprehensive hyperparameter
699 combinations not considered in this study. Such advancements could significantly improve the
700 predictive capabilities beyond the results demonstrated in this study.

701 **5 Conclusions**

702 This study explored optimal methods for regionalizing Intensity-Duration-Frequency (IDF)
703 curves across mainland China, utilizing long-term hourly precipitation data from 2363
704 observation stations across mainland China, as well as daily gridded precipitation data from the
705 CHM_PRE. To achieve this objective, we evaluated and compared the performance of five
706 traditional interpolation methods (site-observation-based) and five emerging machine learning
707 methods (gridded-precipitation-based). Four representative rainfall intensities were selected as
708 examples: 1-hour, 5-year (SDSR), 1-hour, 100-year (SDLR), 24-hour, 5-year (LDSR), and 24-hour,
709 100-year (LDLR). The following key findings were obtained:

- 710 (1) Among all interpolation methods evaluated in this study, Kriging methods incorporating
711 annual precipitation as a covariate (KED_AP) significantly improved prediction accuracy
712 compared with the Ordinary Kriging(OK), whereas including elevation as a covariate did
713 not improve the accuracy. Among all machine learning methods evaluated in this study,
714 Gradient Boosting (GB) demonstrated the best and most robust performance.
- 715 (2) For the best-performing interpolation and machine learning methods across the four
716 rainfall intensities selected, both methods demonstrated comparable and robust
717 performance across various cases, whether for short or long durations and return periods.
718 Specifically, the KED_AP method achieved a KGE of over 0.96 for SDSR and LDSR, and over
719 0.84 for SDLR and LDLR. Similarly, the GB method achieved a KGE of over 0.94 for SDSR
720 and LDSR, and over 0.87 for SDLR and LDLR. Machine learning methods tend to perform

721 better as the duration and return period increase. Besides, shorter durations and larger
722 return periods were associated with increased prediction difficulty and reduced accuracy
723 for both methods, as evidenced by the KGE values for LDSR (KED_AP = 0.96; GB = 0.96)
724 and SDLR (KED_AP = 0.84; GB = 0.87).

725 (3) IDF curves typically require precipitation data with sub-daily temporal resolution (e.g.,
726 hourly) when using traditional interpolation methods. In contrast, machine learning
727 methods, utilizing only daily precipitation data and not previously applied to IDF curves
728 regionalization in mainland China, demonstrated comparable accuracy to traditional
729 methods and even outperformed them for long durations and large return periods.
730 Additionally, machine learning offers distinct advantages, such as only requiring coarser
731 temporal resolution precipitation data, being independent of station setup conditions,
732 and possessing strong adaptability for future updates. These advantages suggest that
733 machine learning has the potential to replace interpolation methods to produce future
734 IDF curves.

735 With the optimal regionalization method for each specific duration and return period, the
736 study ultimately produced a regionalized IDF curves product for mainland China with a spatial
737 resolution of $0.1^\circ \times 0.1^\circ$. The durations considered ranged from 1 to 24 hours, along with 36, 48,
738 and 72 hours, and the return periods included 2, 5, 10, 20, 50, 100, 200, 500, and 1000 years.
739 Although with lower accuracy, another IDF curves product for mainland China with a spatial
740 resolution of $0.5^\circ \times 0.5^\circ$ based on CGDPP, is also provided as an alternative. To resolve potential
741 crossing phenomena within the generated curves, a bivariate isotonic regression was applied as
742 post-processing, ensuring physical consistency with negligible impact on accuracy (Dykstra, 1983).
743 We also provided accuracy metrics and the name of the optimal regionalization methods for each
744 IDF case (Tables S31, S36, S41–S48). Taking the city of Beijing (approximately $16,000 \text{ km}^2$) as an
745 example, the existing hourly gauges provide about 20 station-level IDF curves. Our 0.1° dataset
746 increases this to about 170 available IDF curves, while the 0.5° dataset provides about 8 curves.
747 The 0.1° dataset is recommended for local-scale applications requiring high spatial detail, for
748 urban drainage design and small catchment flood modeling for example, while the 0.5° dataset
749 is appropriate for regional-scale hydrologic modeling where computational efficiency and areal-
750 average representation are prioritized.

751 In summary, this study advances regionalization of IDF curves with several distinct
752 contributions that enhance both methodological understanding and practical application. First,
753 through a systematic comparison of traditional interpolation and emerging machine learning
754 methods, this research highlights the relative strengths of each method, enabling the
755 identification of optimal regionalization strategies tailored to specific duration and return period
756 combinations. Second, the study demonstrates that machine learning methods, using readily
757 available gridded daily precipitation data, can estimate sub-daily IDF curves with accuracy
758 comparable to traditional interpolation methods using hourly data. This finding is particularly
759 significant for regions where high-temporal-resolution data are scarce, as it expands the potential
760 for effective IDF regionalization in data-limited environments. Third, by generating a regionalized
761 dataset on IDF curves for mainland China where such datasets were previously scarce, this study
762 has generated a valuable resource for flood risk assessment and infrastructure planning, while
763 providing a valuable methodological framework for similar studies in other regions.

765 **Table A1. List of important acronyms and their definitions**

Acronym	Definition
ET	Extremely Randomized Trees
GB	Gradient Boosting
GEV	Generalized Extreme Value
IDF	Intensity-Duration-Frequency
IDW	Inverse Distance Weighting
KED	Kriging with External Drift
KED_AP	Kriging with External Drift using Mean Annual Precipitation
KED_DEM	Kriging with External Drift using Elevation
KGE	Kling-Gupta Efficiency
LDLR	Long Duration-Large Return Period
LDSR	Long Duration-Small Return Period
LR	Linear Regression
ML	Machine Learning
MLP	Multilayer Perceptron
NE	Northeastern Monsoon Region
NSE	Nash-Sutcliffe Efficiency
NW	Northwestern Arid Region
OK	Ordinary Kriging
PBIAS	Percent Bias
RF	Random Forest
RMSE	Root Mean Square Error
SDLR	Short Duration-Large Return Period
SDSR	Short Duration-Small Return Period
SE	Southeastern Monsoon Region
SW	Southwestern Tibetan Plateau Region

766

767

768 **Author contribution**

769 YJ developed the methodology, performed the data analysis, created the visualizations, and
770 wrote the original draft; WW conceptualized and designed the study, provided the data and
771 initial scripts, acquired the funding, and supervised the project; WW and YJ revised the original
772 draft; All authors commented, reviewed and edited the manuscript.

773

774 **Acknowledgments**

775 This work was supported by the National Natural Science Foundation of China (No.42307424),
776 and State Key Laboratory of Earth Surface Processes and Resource Ecology (2023-KF-10). We
777 thank Zeng He for helping prepare the scripts. The authors utilized the AI model DeepSeek-R1
778 (Guo et al., 2025) to enhance the language of this manuscript. The tool was employed solely for
779 improving grammar, syntax, and word choice. The authors reviewed and revised all AI-generated
780 suggestions to ensure the scientific accuracy of the content.

781

782 **Data Availability Statement**

783 Regionalized Intensity-Duration-Frequency curves in mainland China (0.1°/0.5°) are available at
784 <https://data.tpdac.ac.cn/en/disallow/ba9c23d1-cdbf-471c-90d5-b413d2a9f86a>.

785

786 **Competing interests**

787 The authors declare that they have no known competing financial interests or personal
788 relationships that could have appeared to influence the work reported in this paper.

789

790 **References**

- 791 Al Kajbaf, A., Bensi, M., & Brubaker, K. L. (2022). Temporal downscaling of precipitation from
792 climate model projections using machine learning. *Stochastic Environmental Research and*
793 *Risk Assessment*, 36(8), 2173–2194. <https://doi.org/10.1007/s00477-022-02259-2>
- 794 Al Kajbaf, A., Bensi, M., & Brubaker, K. L. (2023). Drivers of uncertainty in precipitation
795 frequency under current and future climate – application to Maryland, USA. *Journal of*
796 *Hydrology*, 617, 128775. <https://doi.org/10.1016/j.jhydrol.2022.128775>
- 797 Armstrong, M. (1998). *Basic linear geostatistics*. Berlin, Heidelberg: Springer-Verlag.
798 <https://doi.org/10.1007/978-3-642-58727-6>
- 799 Benestad, R. E., Lutz, J., Dyrddal, A. V., Haugen, J. E., Parding, K. M., & Dobler, A. (2021). Testing a
800 simple formula for calculating approximate intensity-duration-frequency curves.
801 *Environmental Research Letters*, 16(4), 044009. [https://doi.org/10.1088/1748-](https://doi.org/10.1088/1748-9326/abd4ab)
802 [9326/abd4ab](https://doi.org/10.1088/1748-9326/abd4ab)
- 803 Berndt, C., & Haberlandt, U. (2018). Spatial interpolation of climate variables in Northern
804 Germany—Influence of temporal resolution and network density. *Journal of Hydrology:*
805 *Regional Studies*, 15, 184–202. <https://doi.org/10.1016/j.ejrh.2018.02.002>
- 806 Berndt, C., Rabiei, E., & Haberlandt, U. (2014). Geostatistical merging of rain gauge and radar
807 data for high temporal resolutions and various station density scenarios. *Journal of*
808 *Hydrology*, 508, 88–101. <https://doi.org/10.1016/j.jhydrol.2013.10.028>
- 809 Breiman, L. (2001). Random forests. *Machine Learning*, 45, 5–32.
810 <https://doi.org/10.1023/A:1010933404324>
- 811 Courty, L. G., Wilby, R. L., Hillier, J. K., & Slater, L. J. (2019). Intensity-duration-frequency curves
812 at the global scale. *Environmental Research Letters*, 14(8), 084045.
813 <https://doi.org/10.1088/1748-9326/ab370a>
- 814 Donat, M. G., Alexander, L. V., Herold, N., & Dittus, A. J. (2016). Temperature and precipitation
815 extremes in century-long gridded observations, reanalyses, and atmospheric model
816 simulations. *Journal of Geophysical Research: Atmospheres*, 121, 11,174–11,189.
817 <https://doi.org/10.1002/2016JD025480>

818 Dykstra, R. L. (1983). An algorithm for restricted least squares regression. *Journal of the*
819 *American Statistical Association*, 78(384), 837 – 842.
820 <https://doi.org/10.1080/01621459.1983.10477029>

821 Förster, K., & Thiele, L. B. (2020). Variations in sub-daily precipitation at centennial scale. *npj*
822 *Climate and Atmospheric Science*, 3(1), 13. <https://doi.org/10.1038/s41612-020-0117-1>

823 Friedman, J. H. (2001). Greedy Function Approximation: A Gradient Boosting Machine. *The*
824 *Annals of Statistics*, 29(5), 1189–1232. <http://www.jstor.org/stable/2699986>

825 Funk, C., Peterson, P., Landsfeld, M., Pedreros, D., Verdin, J., Shukla, S., Husak, G., Rowland, J.,
826 Harrison, L., Hoell, A., & Michaelsen, J. (2015). The climate hazards infrared precipitation
827 with stations—A new environmental record for monitoring extremes. *Scientific Data*, 2(1),
828 150066. <https://doi.org/10.1038/sdata.2015.66>

829 Gao, L., Huang, J., Chen, X., Chen, Y., & Liu, M. (2017). Risk of Extreme Precipitation under
830 Nonstationarity Conditions during the Second Flood Season in the Southeastern Coastal
831 Region of China. *Journal of Hydrometeorology*, 18(3), 669–681.
832 <https://doi.org/10.1175/JHM-D-16-0119.1>

833 Gaur, A., Schardong, A., & Simonovic, S. P. (2020). Gridded Extreme Precipitation Intensity-
834 Duration-Frequency Estimates for the Canadian Landmass. *JOURNAL OF HYDROLOGIC*
835 *ENGINEERING*, 25(6), 05020006. [https://doi.org/10.1061/\(ASCE\)HE.1943-5584.0001924](https://doi.org/10.1061/(ASCE)HE.1943-5584.0001924)

836 Geurts, P., Ernst, D., & Wehenkel, L. (2006). Extremely randomized trees. *Machine learning*, 63,
837 3-42. <https://doi.org/10.1007/s10994-006-6226-1>

838 Ghebreyesus, D. T., & Sharif, H. O. (2021). Development and Assessment of High-Resolution
839 Radar-Based Precipitation Intensity-Duration-Curve (IDF) Curves for the State of Texas.
840 *Remote Sensing*, 13(15), 2890. <https://doi.org/10.3390/rs13152890>

841 Guo, D., Yang, D., Zhang, H., Song, J., Wang, P., Zhu, Q., Xu, R., Zhang, R., Ma, S., Bi, X., Zhang, X.,
842 Yu, X., Wu, Y., Wu, Z. F., Gou, Z., Shao, Z., Li, Z., Gao, Z., Liu, A., ... Zhang, Z. (2025).
843 DeepSeek-R1 incentivizes reasoning in LLMs through reinforcement learning. *Nature*,
844 645(8081), 633–638. <https://doi.org/10.1038/s41586-025-09422-z>

845 Halbert, K., Nguyen, C. C., Payrastre, O., & Gaume, E. (2016). Reducing uncertainty in flood
846 frequency analyses: A comparison of local and regional approaches involving information

847 on extreme historical floods. *Journal of Hydrology*, 541, 90–98.
848 <https://doi.org/10.1016/j.jhydrol.2016.01.017>

849 Han, J., Miao, C., Gou, J., Zheng, H., Zhang, Q., & Guo, X. (2023). A new daily gridded
850 precipitation dataset for the Chinese mainland based on gauge observations. *Earth System
851 Science Data*, 15(7), 3147–3161. <https://doi.org/10.5194/essd-15-3147-2023>

852 Haruna, A., Blanchet, J., & Favre, A. (2024). Estimation of Intensity-Duration-Area-Frequency
853 Relationships Based on the Full Range of Non-Zero Precipitation From Radar-Reanalysis
854 Data. *Water Resources Research*, 60(2), e2023WR035902.
855 <https://doi.org/10.1029/2023WR035902>

856 Hinton, G. E. (1989). Connectionist learning procedures. *Artificial Intelligence*, 40(1-3), 185-234.
857 [https://doi.org/10.1016/0004-3702\(89\)90049-0](https://doi.org/10.1016/0004-3702(89)90049-0)

858 Hosking, J.R.M. (1990). L-moments: Analysis and estimation of distributions using linear
859 combinations of order statistics. *Journal of the Royal Statistical Society: Series B
860 (Methodological)*, 52(1), 105–124. <https://doi.org/10.1111/j.2517-6161.1990.tb01775.x>

861 Jenkinson, A. F. (1955). The frequency distribution of the annual maximum (or minimum) values
862 of meteorological elements. *Quarterly Journal of the Royal Meteorological Society*,
863 81(348), 158–171. <https://doi.org/10.1002/qj.49708134804>

864 Koutsoyiannis, D., Kozonis, D., & Manetas, A. (1998). A mathematical framework for studying
865 rainfall intensity-duration-frequency relationships. *Journal of Hydrology*, 206(1–2), 118–
866 135. [https://doi.org/10.1016/S0022-1694\(98\)00097-3](https://doi.org/10.1016/S0022-1694(98)00097-3)

867 Lanciotti, S., Ridolfi, E., Russo, F., & Napolitano, F. (2022). Intensity-Duration-Frequency Curves in
868 a Data-Rich Era: A Review. *Water*, 14(22), 3705. <https://doi.org/10.3390/w14223705>

869 Li, J., & Heap, A. D. (2008). *A review of spatial interpolation methods for environmental
870 scientists*. Canberra, Australia: Geoscience Australia.

871 Li, J., Yu, R., Yuan, W., & Chen, H. (2011). Changes in Duration-Related Characteristics of Late-
872 Summer Precipitation over Eastern China in the Past 40 Years. *Journal of Climate*, 24(21),
873 5683–5690. <https://doi.org/10.1175/JCLI-D-11-00009.1>

874 Lundberg, S. M., & Lee, S.-I. (2017). A unified approach to interpreting model predictions. In I.
875 Guyon, U. V. Luxburg, S. Bengio, H. Wallach, R. Fergus, S. Vishwanathan, & R. Garnett

876 (Eds.), *Advances in Neural Information Processing Systems* (Vol. 30, pp. 4765–4774).

877 Maggioni, V., & Massari, C. (Eds.). (2019). *Extreme Hydroclimatic Events and Multivariate*
878 *Hazards in a Changing Environment*. Amsterdam, Netherlands: Elsevier.
879 <https://doi.org/10.1016/B978-0-12-814899-0.00016-X>

880 Miao, C., Immerzeel, W. W., Xu, B., Yang, K., Duan, Q., & Li, X. (2024). Understanding the Asian
881 water tower requires a redesigned precipitation observation strategy. *Proceedings of the*
882 *National Academy of Sciences*, 121(23), e2403557121.
883 <https://doi.org/10.1073/pnas.2403557121>

884 Mínguez, R., & Herrera, S. (2023). Spatial extreme model for rainfall depth: Application to the
885 estimation of IDF curves in the Basque country. *Stochastic Environmental Research and*
886 *Risk Assessment*, 37(8), 3117–3148. <https://doi.org/10.1007/s00477-023-02440-1>

887 Mockus, J., Tiesis, V., & Žilinskas, A. (1978). The application of Bayesian methods for seeking the
888 extremum. In L. C. W. Dixon & G. P. Szegő (Eds.), *Towards global optimisation* (Vol. 2, pp.
889 117–129).

890 Nguyen, V.-T.-V., Nguyen, T.-D., & Ashkar, F. (2002). Regional frequency analysis of extreme
891 rainfalls. *Water Science and Technology*, 45(2), 75–81.
892 <https://doi.org/10.2166/wst.2002.0030>

893 Noor, M., Ismail, T., Shahid, S., Asaduzzaman, M., & Dewan, A. (2021). Evaluating intensity-
894 duration-frequency (IDF) curves of satellite-based precipitation datasets in Peninsular
895 Malaysia. *Atmospheric Research*, 248, 105203.
896 <https://doi.org/10.1016/j.atmosres.2020.105203>

897 Papalexiou, S. M. (2018). Unified theory for stochastic modelling of hydroclimatic processes:
898 Preserving marginal distributions, correlation structures, and intermittency. *Advances in*
899 *Water Resources*, 115, 234–252. <https://doi.org/10.1016/j.advwatres.2018.02.013>

900 Papalexiou, S. M., & Koutsoyiannis, D. (2013). Battle of extreme value distributions: A global
901 survey on extreme daily rainfall. *Water Resources Research*, 49(1), 187–201.
902 <https://doi.org/10.1029/2012WR012557>

903 Parding, K. M., Benestad, R. E., Dyrørdal, A. V., & Lutz, J. (2023). A principal-component-based
904 strategy for regionalisation of precipitation intensity-duration-frequency (IDF) statistics.

905 *HYDROLOGY AND EARTH SYSTEM SCIENCES*, 27(20), 3719–3732.
906 <https://doi.org/10.5194/hess-27-3719-2023>

907 Pebesma, E. J. (2004). Multivariable geostatistics in S: the gstat package. *Computers &*
908 *geosciences*, 30(7), 683-691. <https://doi.org/10.1016/j.cageo.2004.03.012>

909 Pedregosa, F., Varoquaux, G., Gramfort, A., Michel, V., Thirion, B., Grisel, O., ... & Duchesnay, É.
910 (2011). Scikit-learn: Machine learning in Python. *Journal of machine Learning research*, 12,
911 2825-2830. <https://doi.org/10.48550/arXiv.1201.0490>

912 Ren, Z., Sang, Y.-F., Cui, P., Chen, F., & Chen, D. (2025). A dataset of gridded precipitation
913 intensity-duration-frequency curves in Qinghai-Tibet Plateau. *Scientific Data*, 12(1), 3.
914 <https://doi.org/10.1038/s41597-024-04362-1>

915 Sangüesa, C., Pizarro, R., Ingram, B., Ibáñez, A., Rivera, D., García-Chevesich, P., Pino, J., Pérez, F.,
916 Balocchi, F., & Peña, F. (2023). Comparing Methods for the Regionalization of
917 Intensity–Duration–Frequency (IDF) Curve Parameters in Sparsely-Gauged and Ungauged
918 Areas of Central Chile. *Hydrology*, 10(9), 179. <https://doi.org/10.3390/hydrology10090179>

919 Schilcher, U., Brandner, G., & Bettstetter, C. (2017). Quantifying inhomogeneity of spatial point
920 patterns. *Computer Networks*, 115, 65–81. <https://doi.org/10.1016/j.comnet.2016.12.018>

921 Schlef, K. E., Kunkel, K. E., Brown, C., Demissie, Y., Lettenmaier, D. P., Wagner, A., Wigmosta, M.
922 S., Karl, T. R., Easterling, D. R., Wang, K. J., François, B., & Yan, E. (2023). Incorporating non-
923 stationarity from climate change into rainfall frequency and intensity-duration-frequency
924 (IDF) curves. *Journal of Hydrology*, 616, 128757.
925 <https://doi.org/10.1016/j.jhydrol.2022.128757>

926 Shao, D., & Liu, G. (2018). Up-to-date urban rainstorm intensity formulas considering spatial
927 diversity in China. *Environmental Earth Sciences*, 77, 541. [https://doi.org/10.1007/s12665-](https://doi.org/10.1007/s12665-018-7718-6)
928 [018-7718-6](https://doi.org/10.1007/s12665-018-7718-6)

929 Shehu, B., Willems, W., Stockel, H., Thiele, L.-B., & Haberlandt, U. (2023). Regionalisation of
930 rainfall depth–duration–frequency curves with different data types in Germany. *Hydrology*
931 *and Earth System Sciences*, 27(5), 1109–1132. <https://doi.org/10.5194/hess-27-1109-2023>

932 Shen, Y., Mingnong, F., Hongzheng, Z., & Feng, G. (2010). Interpolation Methods of China Daily
933 Precipitation Data. *Journal of Applied Meteorological Science*, 21(3), 279-286.

934 <https://doi.org/10.3969/j.issn.1001-7313.2010.03.003>.

935 Simonović, S. P. (2012). *Floods in a changing climate: Risk management*. Cambridge, UK:
936 Cambridge University Press.

937 Skliris, N., Zika, J. D., Nurser, G., Josey, S. A., & Marsh, R. (2016). Global water cycle amplifying at
938 less than the Clausius-Clapeyron rate. *Scientific Reports*, 6, 38752.
939 <https://doi.org/10.1038/srep38752>

940 Snoek, J., Larochelle, H., & Adams, R. P. (2012). Practical Bayesian optimization of machine
941 learning algorithms. In F. Pereira, C. J. C. Burges, L. Bottou, & K. Q. Weinberger (Eds.),
942 *Advances in Neural Information Processing Systems* (Vol. 25, pp. 2951–2959).

943 Szolgay, J., Parajka, J., Kohnová, S., & Hlavčová, K. (2009). Comparison of mapping approaches of
944 design annual maximum daily precipitation. *Atmospheric Research*, 92(3), 289–307.
945 <https://doi.org/10.1016/j.atmosres.2009.01.009>

946 Verworn, A., & Haberlandt, U. (2011). Spatial interpolation of hourly rainfall – effect of
947 additional information, variogram inference and storm properties. *Hydrology and Earth
948 System Sciences*, 15(2), 569–584. <https://doi.org/10.5194/hess-15-569-2011>

949 Vinod, D., & Mahesha, A. (2024). Modeling nonstationary intensity-duration-frequency curves
950 for urban areas of India under changing climate. *Urban Climate*, 56, 102065.
951 <https://doi.org/10.1016/j.uclim.2024.102065>

952 Wambura, F. J. (2024). Using reanalysis precipitation data for developing intensity-duration-
953 frequency curves in a poorly gauged city. *Journal of Hydrology: Regional Studies*, 56,
954 102005. <https://doi.org/10.1016/j.ejrh.2024.102005>

955 Wang, W., Yin, S., Xie, Y., Nearing, M. A., Wang, W., Yin, S., Xie, Y., & Nearing, M. A. (2019).
956 Minimum Inter-Event Times for Rainfall in the Eastern Monsoon Region of China.
957 *Transactions of the ASABE*, 62(1), 9–18. <https://doi.org/10.13031/trans.12878>

958 Wang, W., Yin, S., He, Z., Chen, D., Wang, H., & Klik, A. (2023). Projections of rainfall erosivity in
959 climate change scenarios for mainland China. *CATENA*, 232, 107391.
960 <https://doi.org/10.1016/j.catena.2023.107391>

961 Westra, S., Fowler, H. J., Evans, J. P., Alexander, L. V., Berg, P., Johnson, F., Kendon, E. J.,
962 Lenderink, G., & Roberts, N. M. (2014). Future changes to the intensity and frequency of

963 short-duration extreme rainfall. *Reviews of Geophysics*, 52(3), 522–555.
964 <https://doi.org/10.1002/2014RG000464>

965 Yin, S., Wang, Z., Zhu, Z., Zou, X., & Wang, W. (2018). Using Kriging with a heterogeneous
966 measurement error to improve the accuracy of extreme precipitation return level
967 estimation. *Journal of Hydrology*, 562, 518–529.
968 <https://doi.org/10.1016/j.jhydrol.2018.04.064>

969 Zhang, B., Wang, S., Moradkhani, H., Slater, L., & Liu, J. (2022). A Vine Copula-Based Ensemble
970 Projection of Precipitation Intensity–Duration–Frequency Curves at Sub-Daily to Multi-Day
971 Time Scales. *Water Resources Research*, 58(11), e2022WR032658.
972 <https://doi.org/10.1029/2022WR032658>

973 Zhao, S. (1983). A new scheme for comprehensive physical regionalization in China. *Acta*
974 *Geographica Sinica*, 38(1), 1–10.

975 Zou, W., Yin, S., & Wang, W. (2021). Spatial interpolation of the extreme hourly precipitation at
976 different return levels in the Haihe River basin. *Journal of Hydrology*, 598, 126273.
977 <https://doi.org/10.1016/j.jhydrol.2021.126273>
978