

Assessing the Impact of the Human Development Index on Historical Trends in the INFERNO Fire Model

João C. M. Teixeira^{1,2,4}, Chantelle Burton¹, Douglas I. Kelley³, Gerd A. Folberth¹, Fiona M. O'Connor^{1,2}, Richard A. Betts^{1,2}, and Apostolos Voulgarakis^{4,5}

¹Met Office, Fitzroy Road, EX1 3PB, Exeter, UK

²Global Systems Institute, Department of Mathematics & Statistics, University of Exeter, EX4 4QE, UK

³UK Centre for Ecology and Hydrology, Wallingford OX10 8BB, U.K

⁴Leverhulme Centre for Wildfires, Environment and Society, Department of Physics, Imperial College London, London, UK

⁵School of Environmental Engineering, Technical University of Crete, Chania, Greece

Correspondence: João C. M. Teixeira (joao.teixeira@metoffice.gov.uk)

Abstract. Earth System Models (ESMs) capture long-term historical trends in burnt area, but they struggle to reproduce the pronounced decline observed over the past two decades, largely because anthropogenic fire suppression and related socio-economic influences are not adequately represented. Key factors such as agricultural expansion, land-use changes, fire management policies, and landscape fragmentation have all contributed to reduced fire activity, especially in tropical savannas, but these are not adequately captured in the fire model formulations that underpin most ESMs.

This study investigates whether the observed downward trend in global burnt area during the period 1998–2016 can be better represented in the JULES-INFERNO fire model by incorporating a simplified representation of direct human impacts on fire. Specifically, we focus on the Human Development Index (HDI), which reflects socio-economic development and, in turn, influences fire suppression efforts. By implementing globally uniform assumption that anthropogenic fire ignitions and suppression decline linearly with HDI into INFERNO, we aim to improve the representation of fire ignition and suppression dynamics.

Including HDI-based socio-economic factors substantially reduces regional biases in annual burned area. In Temperate North America, for example, model bias decreases from +735.57 % to +44.46 %, with similarly large reductions in Central America, Southern Hemisphere South America, Europe, and the Middle East. HDI also improves the representation of burned area trends in eight of the 14 GFED4s regions that exhibit significant negative trends in observations. However, correcting large positive regional biases removes compensating errors in the original configuration, resulting in a stronger global negative bias, shifting from -34.35 Mha to approximately -111 Mha.

While HDI enhances regional performance and better captures recent downward trends in some areas, it introduces trade-offs, including dampened interannual variability and underestimation of medium to large fire events. Overall, incorporating HDI offers a computationally efficient way to represent socio-economic influences within INFERNO. Nevertheless, the linear and spatially uniform implementation remains a first-order approximation and cannot fully capture the complex regional, cultural, and policy-driven dynamics shaping human–fire interactions.

1 Introduction

Globally, burnt area trends are influenced by complex interactions between climate change, human activities, and natural ecosystem processes, resulting in large climate change and variability over the past few decades. However, the long-term trend (e.g., 1997 - 2016) has shown an overall decline in global burnt area, especially driven by changes in burnt area of African savannas and grasslands (decline of 1.27 % per year (Andela et al., 2017)). The trend, shown in Figure 1 b), is attributed to changes in land use, particularly agricultural expansion and intensification in savanna and grassland regions, which reduces the availability of fuel for fires (Riley et al., 2019; Andela et al., 2017).

In addition, fire is also actively used as a land management tool, for example to clear land for agriculture, manage vegetation, and maintain pasture systems, particularly in tropical and subtropical landscapes.

Figure 1 illustrates the broad-scale spatial patterns and long-term trends of burnt area, fire weather, and the Human Development Index (HDI).

Climate is a key factor that also influences fire activity. Rising temperatures are leading to longer fire seasons, particularly in temperate and boreal regions (Sullivan et al., 2022; Jones et al., 2024). Additionally, reduced rainfall during the critical phase of fire seasons increases the likelihood of large-scale fires, while increased rainfall during certain times of the year can promote vegetation growth, providing more fuel for future fires. However, while climate strongly influences inter-annual variability and is increasingly important for long-term trends, particularly in temperate and boreal regions, human activity—such as land use change, agricultural expansion, and active fire suppression—has been the dominant driver of long-term declines in global burnt area, especially in tropical savannas (Riley et al., 2019; Andela et al., 2017; Marlon et al., 2008).

Furthermore, human population density and prosperity can significantly impact burnt areas. (Andela et al., 2017) show that due to population growth, socioeconomic development, and the increased demand for agricultural products in regional and global markets, there has been a shift towards more capital-intensive agriculture, resulting in fewer and smaller fires. These factors have a predictable impact on the use of fire, with a strong inverse correlation between the area burned and economic development.

In Coupled Model Intercomparison Project frameworks, fire processes are generally represented through simplified, global-scale schemes that link burned area to vegetation, fuel availability, and climate conditions. Most models include ignition from lightning and human sources, with fire spread and, flammability, and fuel continuity. These schemes capture broad spatio-temporal patterns of fire but tend to not resolve fine-scale processes such as local suppression, management practices, or sub-grid fuel heterogeneity (Yue et al., 2014; Rabin et al., 2017).

The study by Li et al. (2024) shows that the Earth System Models (ESMs) used to provide state-of-the-art climate projections for Phase 6 of the Coupled Model Intercomparison Project (CMIP6; Eyring et al. (2016)) do not reproduce the observed decline in global burned area and fire carbon emissions over the past two decades. They identify the primary cause of this discrepancy as an underestimation of anthropogenic fire suppression in fire-enabled models (mathematical representation of a real-world system). This finding is consistent with the data-driven analysis of Forkel et al. (2019), who quantitatively disentangled the

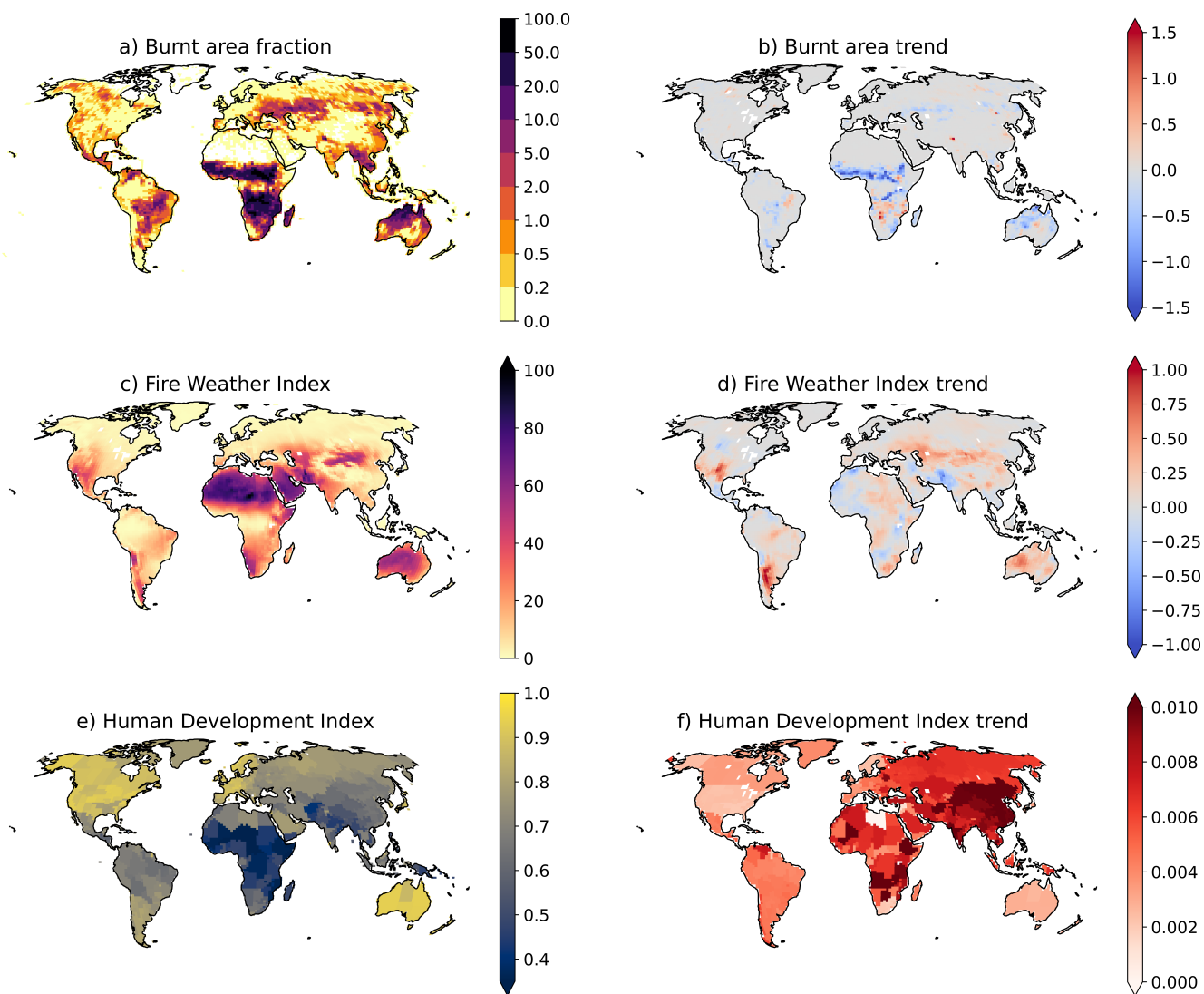


Figure 1. Global distribution of a) Burnt area fraction from the Global Fire Emission Database version 4 (GFED4s) (%) total annual average (1997 - 2016) and b) respective trend ($\% \text{ year}^{-1}$); c) Fire Weather Index (FWI) average across 1997 - 2016 and d) respective trend (year^{-1}); e) Human Development Index (HDI) average across 1997 - 2016 and f) respective trend (year^{-1}). The FWI maps represent long-term climatological averages intended to illustrate broad-scale spatial patterns and trends in fire-conducive conditions.

relative contributions of climate and human drivers of fire activity and demonstrated there is a strong sensitivity to climate variables, vegetation properties, and to socio-economic variables.

Moreover, key human-driven factors—such as agricultural expansion, land-use change, fire management policies, and landscape fragmentation, have substantially reduced fire activity, particularly in tropical savannas (Andela et al., 2017; Forkel et al., 2019). However, most CMIP6 models do not adequately represent these suppression mechanisms, leading to an overestimation of burned area trends and associated fire-related carbon emissions.

Most global fire models, such as JULES-INFERNO (Mangeon et al., 2016; Burton et al., 2019), apply simplistic representations of human ignitions, generally specified as a function of population density, increasing up to a threshold value after which there are no additional ignitions with an increasing population (Rabin et al., 2017; Teckentrup et al., 2019; Ford et al., 2021). This approach does not account for the impacts of socio-economic factors on anthropogenic fire ignitions or suppression. It also does not capture how humans influence regional variation in contemporary burning practices.

This misrepresentation has important implications for climate projections and carbon cycle modelling, and emphasizes the need for better integration of dynamic human influences on fire regimes, such as incorporating evolving socio-economic factors (e.g., population growth, infrastructure expansion, and fire suppression strategies) into fire models (Li et al., 2024).

In current global fire models, changes in fire activity are typically quantified through a separation of climate-driven controls—such as fire weather, fuel moisture, and vegetation productivity—from anthropogenic influences, which are commonly represented using simplified proxies such as population density or land-use change. As shown by (Burton et al., 2024), this structural separation strongly shapes how models attribute historical and future changes in fire activity to climate versus human drivers, and can lead to substantial differences in simulated fire responses when anthropogenic processes are under-represented.

Accurately simulating fire is also essential for reliable climate projections, as fire and climate are coupled through two-way feedbacks. Fire activity influences atmospheric composition, surface albedo, and land–carbon exchanges, while climate change alters fire weather, fuel availability, and ignition conditions. Recent work by (Verjans et al., 2025) highlights the importance of these fire–climate feedbacks, showing that improved fire representations can substantially affect simulated climate responses. This further underscores the need for realistic fire modelling within ESMs.

HDI is a composite measure that combines four key metrics: life expectancy at birth, expected years of schooling, average years of schooling, and Gross National Income (GNI) per capita (Bhanojirao, 1991). These metrics are normalized by their respective maximum values, and the HDI is calculated as the geometric mean of life expectancy, education, and GNI per capita. The HDI has been used in various studies to better understand socio-economic influences on the Earth System, including links between development and environmental pressure (Türe, 2013), the decoupling between human development and resource use or emissions (Hickel, 2020), and the role of socio-economic development pathways in shaping sustainability outcomes at regional and global scales (Roy et al., 2023).

Chuvieco et al. (2021) demonstrates that the HDI is strongly correlated with the inter-annual variability of burned area. Regions with higher HDI show lower variability, largely because increased mechanization and a shift away from agrarian livelihoods reduce the need for fire in agricultural practices. Conversely, areas with lower HDI exhibit greater variability,

90 reflecting continued reliance on fire as a land management tool. Incorporating socio-economic indicators such as HDI into fire models significantly enhances their ability to reproduce observed patterns of variability.

More recently, Perkins et al. (2024) introduced WHAM! (the Wildfire Human Agency Model), a global behavioural, geospatial model designed to represent human fire use and management in a form that can be coupled to dynamic global vegetation models, demonstrated via coupling with JULES-INFERNO. WHAM! is empirically grounded in a global synthesis of anthropogenic fire impacts and seeks to move beyond single-proxy parametrisations by representing underlying behavioural and land-system drivers that shape how people ignite, manage, and suppress fires across regions. As such, WHAM! constitutes a more sophisticated, process-oriented attempt to incorporate anthropogenic drivers in fire modelling. In contrast, the HDI-based implementation in this work is intentionally simplified to remain tractable for ESM applications.

100 Additionally, Li et al. (2013) and Zou et al. (2019) investigated the use of Gross Domestic Product (GDP) to parametrise human influences on fire activity. While these approaches capture broad relationships between economic development and fire occurrence, they offer a limited representation of the underlying socio-economic processes that govern fire ignition and suppression.

Human influences on fire activity have become increasingly pronounced since the late 18th century, driven by industrialization, climate change, land clearance, population growth, the decline of traditional fire management practices, and the emergence of large-scale firefighting and fuel management in the 20th century (Bowman et al., 2020). These trends highlight the need for improved data collection to accurately quantify and model the relationships between human populations and fire activity, including factors such as socio-economic status, historical, cultural, and political legacies. Understanding these interactions is particularly important as the economic and environmental impacts of vegetation fires are expected to intensify under anthropogenic climate change (Sullivan et al., 2022; Jones et al., 2024; Haas et al., 2024), this is increasingly urgent. Additionally, Nikolakis and Roberts (2022) examined how policy learning occurs in wildfire governance, focusing on how wildfire policies in British Columbia have evolved. Their case study shows that policy transfer from similar contexts, particularly Indigenous peoples and their governments, can reshape perceptions of wildfire risk and solutions as we adapt to an uncertain future. Similarly, Pandey et al. (2023) explored the complexity of socio-economic factors on fire and the varied fire management policies worldwide, demonstrating that despite differences, these policies have led to a gradual reduction in fire occurrences and burned area over time.

Several studies indicate that, in developed regions, fire occurrence and impacts are strongly shaped by governance and policy, rather than individual behaviour alone. For example, changes in policy and governance have been shown to influence fire outcomes (Nikolakis and Roberts, 2022), while risk framing and coordination efforts also play an important role (Jacobson et al., 2022). Modelling approaches that emphasise management institutions further highlight the significance of governance in shaping fire dynamics (Ford et al., 2021). Observed reductions in fire numbers or size following prevention and suppression policy shifts provide additional empirical support (Curt and Frejaville, 2018), and studies demonstrate that policy design and its dependence on local stakeholder and contextual factors can determine effectiveness (Carreiras et al., 2014). Finally, institutional capacity and spending have been linked directly to fire outcomes, underscoring the importance of resources and coordination (Mourão and Martinho, 2014).

125 Notably, Curt and Frejaville (2018) found that wildfire policies in Mediterranean France have led to a nearly linear decrease in the number of fires since 1975, though the burnt area has fluctuated more abruptly. Therefore, representing land and fire management policies in global fire models is crucial to building confidence in the modelling frameworks which are used to understand future climate regimes. This, in turn, can underpin decision-making by policy-makers in regards to fire policy in the future.

130 Nonetheless, how socio-economic factors impact on fires is complex and dependent on many factors that are difficult to represent in an ESM context. These factors depend on policies implemented at government level, as well as cultural behaviour which varies widely across the world. As a result, the formulation of Climate and ESMs does not allow for representing these details.

Rather than attempting to capture the full complexity of socio-economic influences on fire, this study explores a simplified, emergent relationship, using HDI as a proxy, that is better suited for large-scale applications. Here, fire suppression is used broadly to represent human actions that limit fire spread, including active containment and land management practices, whose implementation and effectiveness vary regionally with resources and infrastructure. Processes such as fine-scale suppression tactics, planned burns, fuel management, and weather-driven fire extinction are not explicitly represented. We use observed datasets of HDI, burnt area, and Fire Weather Index (FWI) to derive a relationship between HDI and burnt area, and model data to implement this approach in the INFERNO fire model.

140 In Section 2, we explore the relation between HDI and burnt area and describe the INFERNO fire model, the coupling of INFERNO to the latest representation of the land surface model (JULES-ES) as used in the UK's Earth System Model (UKESM1), and how we include HDI into INFERNO's ignition scheme. In Section 3, we evaluate the impact of considering HDI on burnt area, burnt area trends, as well as the impact of external model drivers of burnt area trends. Discussion and conclusions from this work are presented in Section 4, where we focus on novel model results, placing the link between socio-economic factors and fires in context with existing literature. Model limitations and known issues are also highlighted.

2 Methods

2.1 Observed Datasets

2.1.1 The Human Development Index dataset

150 HDI originated from the annual Human Development Reports created by the United Nations Development Programme (UNDP) Human Development Report Office. These reports had the explicit purpose of shifting the focus of development economics from national income accounting to people-centred policies. The aim was to provide a simple composite measure of human development to convince the public, academics, and politicians to evaluate development not only by economic advances but also improvements in human well-being. HDI serves as a crucial metric for assessing the development status of regions globally, and it has been used in several studies to better understand the socio-economics impacts in the Earth System (ES) (Türe, 2013; Hickel, 2020; Roy et al., 2023).

HDI is a composite index (ranging from zero to one) measuring four key metrics (Bhanojirao, 1991):

- life expectancy at birth
- expected years of schooling
- 160 – average years of schooling
- gross national income (GNI) per capita

These metrics are then normalised by their respective maximum value, and HDI is calculated as the geometric mean of life expectancy, education, and GNI per capita, as shown in Eq. 1.

$$HDI = (H_N \cdot E_N \cdot I_N)^{\frac{1}{3}} \quad (1)$$

165 where H_N is the normalised life expectancy, E_N is the normalised arithmetic mean of the two education indices and I_N is the normalised GNI.

The work conducted by Kummu et al. (2018) introduces gridded global datasets for Gross Domestic Product (GDP) and Human Development Index (HDI), covering a 25-year period from 1990 to 2015 with annual frequency and a spatial resolution of 5 arc-minutes. This temporal coverage and high-resolution global scope enable comprehensive analyses of trends, patterns, and changes in HDI across diverse regions and timescales.

170 To produce these datasets, Kummu et al. (2018)s employed a comprehensive approach. For GDP, they utilized both sub-national and national-level data sources. Sub-national GDP data was derived from previous research, while national-level data was sourced from reputable institutions such as the World Bank and the Central Intelligence Agency World Factbook. The HDI dataset was compiled by initially constructing a full national HDI dataset based on data from the Human Development Reports by UNDP. For countries not included in the UNDP reports, independent data sources were utilized, and for missing or outdated data, a methodology involving scaled regional data was adopted.

2.1.2 Burnt Area observation

Global Fire Emission Database version 4 (GFED4s) is a long-running, operationally updated dataset used globally for fire and emissions research. While the core methodological reference (Giglio et al., 2013) describes the development of the GFED4s framework, the dataset itself has been continuously updated to include fire emissions up to recent years (e.g., including 1997–2016).

185 We use data from the GFED4s to understand the relation between HDI and burnt area, as well as to assess the model performance in simulating burnt area. This dataset is provided as a gridded product at a 0.25° resolution. It is derived from a multi-sensor satellite dataset, including satellite data based on active fire detection, and including small fires based on statistical modelling, as detailed in (Randerson et al., 2012).

We apply regions defined in the GFED4s dataset to the modelled data to evaluate the results at a regional level (Figure A1).

2.1.3 Fire Weather Index

The Canadian FWI is a component of the Canadian Forest Fire Danger Rating System (CFFDRS) and provides a numerical rating of fire intensity based solely on weather conditions (temperature, relative humidity, wind speed, and precipitation).

190 Although it was originally developed for Canadian boreal forest conditions, the FWI system is not region-specific and has been successfully applied in diverse ecosystems worldwide, including in Europe, South America, Australia, and parts of Africa and Asia. Its broad adoption stems from its simplicity, weather-based formulation, and scalability (Field et al., 2015).

The work developed by Vitolo et al. (2020) provides an ERA5-based global meteorological wildfire danger dataset based on the Global ECMWF Fire Forecast (GEFF) model and the ERA5 reanalysis.

195 In this dataset the FWI is calculated using meteorological variables (e.g., such as temperature, humidity, precipitation, and wind speed) driven by the ERA5 reanalysis dataset. To accurately represent wildfire conditions at local noon, when fire danger is typically highest, atmospheric fields from ERA5 undergo preprocessing, stitching together hourly forecasts, ensuring that meteorological conditions are representative of 12:00 noon local time around the world.

The GEFF model, calculates the FWI by modelling fuel moisture response to atmospheric forces at different depths. Three 200 fuel moisture levels are used, representing surface fuels, deeper organic material, and compact fuels, each responding at different rates to changes in weather. These moisture codes are then combined to estimate fire behaviour, such as the rate of spread and fire intensity, providing a comprehensive fire danger index.

The GEFF-ERA5 FWI reanalysis dataset, available from 1979 onwards at a spatial resolution of 28 km, was used in this study with monthly averaged values.

205 When comparing datasets (modelled or observed) at different grid resolutions, the higher-resolution dataset is re-gridded to the lowest-resolution grid using a first-order conservative area-weighted re-gridding method.

2.2 Relation between HDI and Burnt Area

As climate is a dominant factor influencing fire activity, it is essential to first account for and remove climate change's influence before exploring the effects of socio-economic factors on burnt area. For this, and considering that the FWI depends solely on 210 the weather variables that drive fire activity, we perform a deweathering procedure at each individual grid cell.

For each grid cell (i, j) , we consider the monthly time series (t) of normalised FWI, $\text{FWI}_{i,j}(t)$, and normalised burnt area, $\text{BA}_{i,j}(t)$, over the study period. A linear regression of the FWI time series is fitted to estimate the local climate-driven trend, $\beta_{i,j}^{\text{FWI}}$, while independently a linear regression of the burnt area time series is used to estimate its intercept, $\alpha_{i,j}^{\text{BA}}$.

The climate-driven component of burnt area is then estimated as

$$215 \quad \widehat{\text{BA}}_{i,j}(t) = \beta_{i,j}^{\text{FWI}} t + \alpha_{i,j}^{\text{BA}} \quad (2)$$

Finally, the climate-driven component is removed from the normalised burnt area time series while preserving the original mean:

$$\text{BA}_{i,j}^*(t) = \text{BA}_{i,j}(t) - \widehat{\text{BA}}_{i,j}(t) + \overline{\text{BA}}_{i,j}, \quad (3)$$

where $\overline{\widehat{BA}}_{i,j} = \frac{1}{T} \sum_{t=1}^T \widehat{BA}_{i,j}(t)$ is the temporal mean of the estimated climate-driven component at each grid cell.

220 This results in a normalised burnt area dataset where the influence of fire-weather variability has been minimized (deweathered), which is then used to analyse how socio-economic factors, through the HDI, impact fire activity.

In this analysis, we make the simplifying assumption that, once the influence of weather is removed, residual variations or trends in burnt area can be primarily attributed to socioeconomic factors. We note, however, that fuel availability and continuity can also be important determinants of fire activity and may influence burnt area trends even after deweathering.

225 To ensure consistency across our burnt area results, we define the burnt area fraction as the proportion of each grid cell that burns in a given year, calculated as the burned area divided by the total cell area (Giglio et al., 2013).

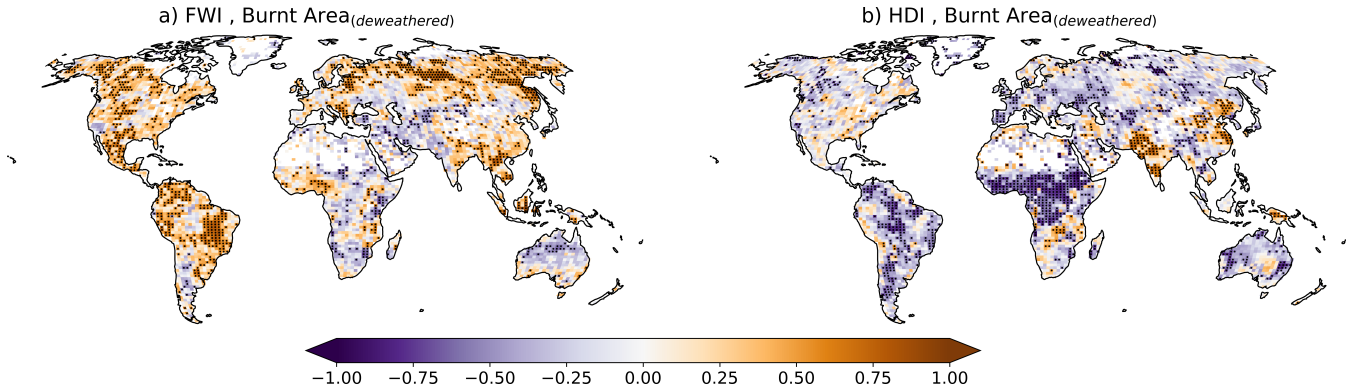


Figure 2. Pearson correlation coefficient between the monthly mean (1997 – 2016) a) FWI and the deweathered Burnt Area (GFED4s), and b) HDI and the deweathered Burnt Area (GFED4s). Stippling indicates grid points where the Pearson correlation is statistically significant at the 5 % level after controlling for the false discovery rate using the Benjamini and Yekutieli (2001) procedure.

Figure 2 shows the spatial correlation coefficient. between FWI and deweathered burnt area (panel a) and between HDI and deweathered burnt area (panel b). While these correlations are spatially variable, over most of the globe the correlation between deweathered burnt area and HDI is not statistically significant. In the regions where the correlation is significant, there is generally a stronger negative correlation between HDI and deweathered burnt area compared to FWI, with notable exceptions in the Indian subcontinent and Eastern China, where a positive correlation is observed. The negative relationship is particularly evident over Eurasia, Southern Africa, and Eastern and Western Australia, where increases in FWI coexist with regional decreases in burnt area that are significantly correlated with HDI (Figures 1 and 2). Nonetheless, the deweathered burnt area shows a strong positive correlation with FWI over the boreal regions of North America and Siberia, as well as the Cerrado ecoregion of South America.

235 Multiple techniques are used to further investigate the relationship between HDI and burnt area. The scatter plot of the deweathered burnt area against HDI in natural space is presented in Figure 3. Due to the inherent stochasticity of burnt area, the results show a wide range of values, spanning small to large relative burnt area fractions for any given HDI. To account for this variability and isolate the effects of HDI on burnt area, we applied a Bayesian Linear Regression method (Klauenberg et al., 2015) to the log-transformed burnt area fraction, denoted by $\log(BA^*)$:

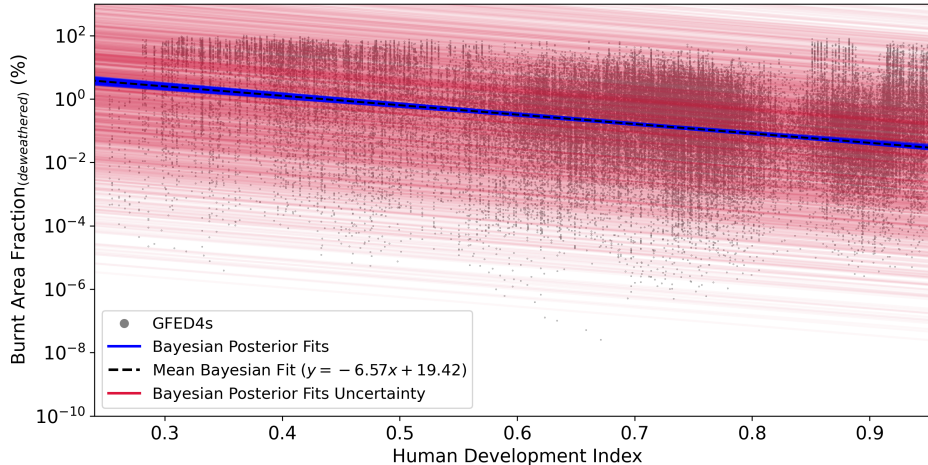


Figure 3. Relation between the monthly mean (1997–2016) HDI and the burnt area fraction (%) for GFED4s. The grey dots represent the scatter plot for all monthly burnt area values at all grid cells as a function of HDI. The Bayesian regression posterior fits are represented by blue solid lines, the mean Bayesian fit is represented by the black dashed line, and the uncertainty from the Bayesian posterior fit residuals is represented by red solid lines. Although the Bayesian regression was performed in log-transformed space, the results are presented here in natural space for interpretability.

$$\log(BA^*) \sim BA_0 + \delta BA \times HDI \quad (4)$$

where BA_0 represents the expected log-burnt area at $HDI = 0$, and δBA represents the change in log-burnt area with HDI. We assumed that $\log(BA^*)$ is normally distributed centred on the linear predictor, with residual standard deviation σ treated as an additional parameter with a half-normal prior. Priors for the regression coefficients were chosen to be weakly informative: $BA_0 \sim \mathcal{N}(0, 25)$ and $\delta BA \sim \text{LogNormal}(0, 10)$, reflecting uncertainty in both the intercept and the HDI effect while constraining δBA to positive values.

The model was fitted using monthly burnt area and HDI data from all grid cells, allowing short-term variability to be smoothed and the longer-term relationship between HDI and burnt area to be quantified. Posterior inference was carried out using the No-U-Turn Sampler (NUTS) implemented in the PyMC Python package (Kelley et al., 2019). Convergence was verified via R-hat values < 1.01 , and 1000 posterior draws were generated to quantify parameter uncertainty. Although the Bayesian regression was performed in log-transformed space, the results are presented here in natural space for interpretability.

To visually represent the uncertainty, 145 posterior draws were plotted as solid blue lines in Figure 3. This number was chosen pragmatically to provide a sufficiently large sample size to capture posterior variability while avoiding excessive clutter in the figure, ensuring a clear representation of both the spread and central tendency of the predicted relationship between HDI and burnt area.

This method shows that the observations show a linear decline in burnt area with increasing HDI, with a posterior mean slope of -6.57 (%).

2.3 INFERNO fire model

In this work we simulate fire using the INFERNO (INteractive Fires and Emissions algoRithm for Natural enviroNments; 260 Mangeon et al. (2016)) fire model. INFERNO uses an approach based on Pechony and Shindell (2009), adapted to allow interactions within an ESM framework. More precisely, INFERNO uses water vapour pressure deficit as one of the main indicators of flammability and an inverse exponential relationship to relate flammability to soil moisture.

$$BA_{PFT} = I_T F_{PFT} \overline{BA_{PFT}} \quad (5)$$

where I_T represents the fire ignitions, including natural and human ignitions as well as fire suppression by humans, F_{PFT} 265 the flammability per PFT dependent on the $1.5 m$ temperature, $1.5 m$ relative humidity and fuel density - as defined in Eq. 4 through 6 from Mangeon et al. (2016) - and $\overline{BA_{PFT}}$ is the average burnt area for each model Plat Functional Type (PFT).

The burnt area, represented in Eq. 5, is the average burnt area per fire for each model PFT. This decouples fire spread from localised effects of wind, weather, or topography, which are not resolved at the coarse spatial scales used in ESMs. Instead, INFERNO relies on PFT-specific flammability and fire occurrence metrics, capturing broad-scale, climate and vegetation- 270 driven fire dynamics. This approach does not explicitly represent sub-grid heterogeneity in topography or local meteorology, which is a limitation that may influence local fire spread.

It should be noted that the recent work by Haas et al. (2022) shows that topography and wind speed have an impact on fire size even when aggregated to a 0.50° grid cell scale. However, the resolution of the model used in this study is coarser - approximately 1.75° .

INFERNO fire ignitions are split into Natural Ignitions (I_N) from cloud to ground lightning flashes and from Human activi- 275 ties (I_A) dependent on population density (PD) as described in Eq. 6. Humans are also responsible for suppressing fires in the model, using a suppression function (f_{NS}) dependent on human population density (Eq. 7) to represent the fraction of fires not suppressed by humans. The total ignitions (I_T) are represented by Eq. 8.

$$I_A = k_{(PD)} PD^\alpha \times (\mathbf{1} - \mathbf{HDI}) \quad (6)$$

$$280 \quad f_{NS} = 7.7 (c_1 + c_2 \times e^{-\omega PD}) \times (\mathbf{1} - \mathbf{HDI}) \quad (7)$$

$$I_T = (I_N + I_A) \frac{f_{NS}}{8.64 \times 10^{10}} \quad (8)$$

where $k_{(PD)} = 6.8 \times PD - 0.6$ is a function that represents the varying anthropogenic influence on ignitions in rural versus urban environments, and the parameter $\alpha = 0.03$ represents the number of potential ignition sources per person per month per km^2 , and HDI represents the Human Development Index.

285 In Equation 7 the fraction of fires that remain unsuppressed at the most populated areas is expressed by c_1 . The maximum number of fires that remain unsuppressed at the distant, unpopulated regions is defined by the sum of c_1 and c_2 , and the rate at which the number of unsuppressed fires decreases with increasing population density is determined by ω . As expressed by Pechony and Shindell (2009), due to the lack of global quantitative data, constant values are selected in a rather heuristic manner: $c_1 = 0.05$, $c_2 = 0.9$, and $\omega = 0.05$. In this way, up to 95 % of fires are assumed to be suppressed in densely populated
290 regions, and 95 % are assumed to remain unsuppressed in unpopulated regions.

Previously, INFERNO only included information on population density. To represent the socio-economic factors impacting fire ignition and suppression, we include a HDI term ($1 - HDI$) in our human ignition and suppression Eq. 6 and 7 (shown in bold). In addition it should be noted that the HDI implementation scales both c_1 and c_2 according to the HDI value of any given grid point.

295 This approach does not directly implement the empirical relationship established from observational data in Section 2.2. Instead, it introduces an HDI dependent formulation based on the assumption that ignition rates decrease and suppression increases with higher HDI. This reflects the hypothesis that more developed regions experience fewer ignitions and greater fire control capacity. We then test whether this indirect implementation can reproduce the observed relationship between burned area and HDI.

300 Despite being a simple representation, it aligns with the few studies found in literature that looked at the impact governmental policies have on prevention of wildfires (e.g., the work by Curt and Frejaville (2018)). Furthermore, although the equations used could be adjusted to provide the best results, we avoid this approach in this first implementation in INFERNO to avoid masking compensating biases that are existent or could arise from this implementation.

In this representation of socio-economic impacts on fire ignition and suppression, we assume that fire ignitions decrease and
305 fire suppression increases for areas with more effort in human development improvements. Moreover, it reduces the impact changes in population density have in areas with high HDI while keeping a dependency on population density changes for areas with low HDI, where policies on land and fire management have a greater role than other human behaviours in controlling ignitions (Nikolakis and Roberts, 2022; Ford et al., 2021; Jacobson et al., 2022; Carreiras et al., 2014; Mourão and Martinho, 2014). The impact of HDI on INFERNO anthropogenic fire ignitions and suppression, represented as I_A (Eq. 6) and f_{NS} (Eq.
310 7) respectively, is depicted in Figure 4.

We obtained HDI data from the gridded global datasets for Gross Domestic Product and Human Development Index (Kummu et al., 2018), which provides HDI data from 1990 to 2015. To cover the full modelled period (1860-2016), the HDI data is linearly ramped from the minimum HDI value of the dataset (0.2) to its value in 1990 for each grid point. The original HDI dataset was spatially interpolated using a nearest-neighbor interpolation method to match the model grid, and was updated at
315 the same frequency as the original dataset - annually.

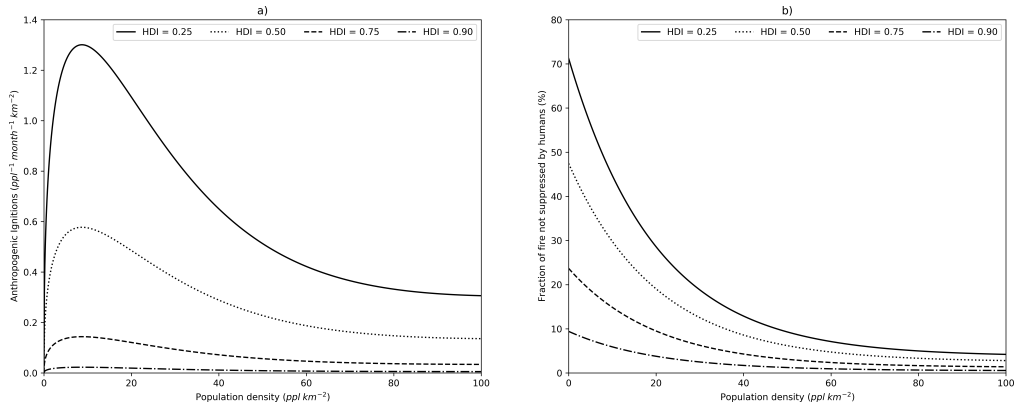


Figure 4. Anthropogenic fire ignitions ($ppl^{-1} month^{-1} km^{-2}$) a) and fraction of fire not suppressed by humans (%) b) as a function of population density ($ppl km^{-2}$) and Human Development Index.

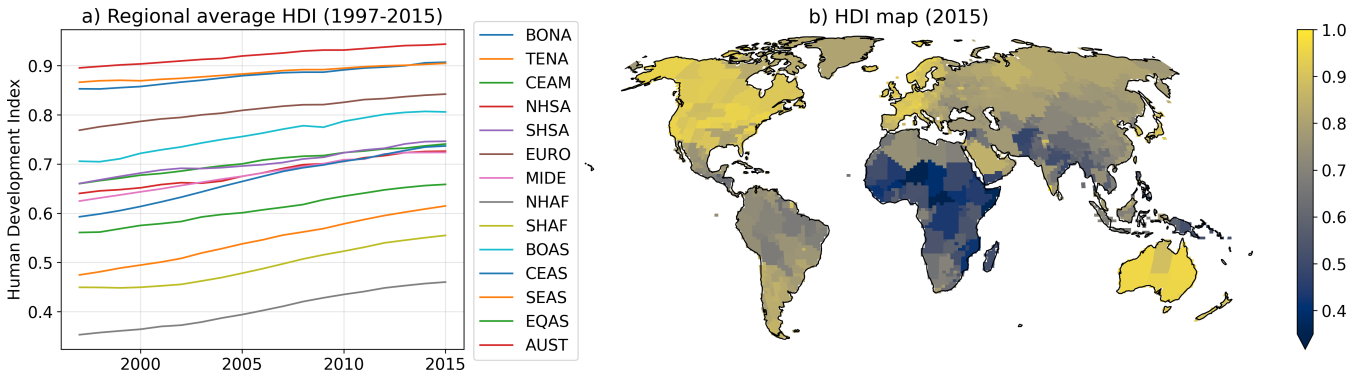


Figure 5. RegridDED HDI as provided to JULES-INFEERNO+HDI. a) regional average for the 1997-2015 period and b) spatial distribution for the year 2015. The regions depicted in a) are described in Figure A1.

2.4 JULES-ES and INFERNO

We use the community land surface model JULES (Joint UK Land Environment Simulator; Clark et al. (2011); Best et al. (2011)) at version 5.7, with the science configuration of the land surface as used in UKESM1 (Sellar et al., 2019a), including 13 PFTs, and dynamic vegetation from TRIFFID (Top-down Representation of Interactive Foliage and Flora Including Dynamics; Cox et al. (2000); Cox (2001)). This ES configuration of JULES is known as JULES-ES (Mathison et al., 2022). JULES simulates surface fluxes of water, energy, as well as vegetation and carbon. Here, we use JULES as a stand-alone offline model run at a spatial resolution of N96 (equivalent to a horizontal resolution of 135 km in the mid-latitudes). The Climate Research Unit - National Centers for Environmental Predictions reanalysis (CRU-NCEP v7) (Harris et al., 2014; Viovy, 2018) atmospheric variables are provided at 6-hourly intervals to drive JULES, including carbon dioxide (CO_2), precipitation, tem-

325 perature, specific humidity, wind, air pressure, and short and long wave radiation. The model runs from 1860–2016 with this forcing. In this work, we analyse the period that overlaps with observations (1997-2015).

We use the fire–vegetation coupling described in Burton et al. (2019, 2020), which incorporates additional carbon cycle feedbacks from litter and vegetation burning and explicitly accounts for fire-induced mortality by plant functional type (PFT). Mortality rates, set to 40 % for trees, 60 % for shrubs, and 100 % for grasses, are based on literature values and the methodology
330 documented in Burton et al. (2019), which includes a comprehensive evaluation of the model performance in representing the evolution of vegetation within the context of JULES-INFERNNO model. This setup differs from that used in Teixeira et al. (2021), which did not include fire–vegetation feedbacks. While simplified, these PFT-specific mortality parameters are intended to capture large-scale patterns of fire-driven vegetation change rather than site-specific fire regimes.

This approach represents a simplification of real-world fire processes. Due to spatial resolution and computational con-
335 straints, global fire modelling frameworks necessarily abstract the diversity and complexity of fire behaviour across ecosystems, and variations in fuel type, structure, and landscape heterogeneity are not fully resolved at the model scale. Despite these limitations, coupling with JULES allows INFERNNO to capture broad spatial and temporal patterns in fuel availability consistent with large-scale vegetation dynamics and meteorological conditions.

Fire ignitions are based on population density data from HYDE 3.2 (Klein Goldewijk et al., 2017); (Goldewijk et al., 2017)
340 and monthly lightning flash climatology from LIS–OTD (Lightning Imaging Sensor–Optical Transient Detector; Cecil (2006)) observations over 1995–2014, regridded from 0.5° resolution to N96 (1.25° latitude × 1.875° longitude). The LIS-OTD climatology provides total lightning flash density (including both intra-cloud and cloud-to-ground flashes); following Christian et al. (2003), total lightning flashes are converted to cloud-to-ground flash density using an empirical partitioning based on observed global relationships between total and cloud-to-ground lightning. After spinning up the model to equilibrium, we
345 complete a full historical simulation from 1860–2019 at N96 and use results from the present day period (1997-2015) for our analysis, which are compared against available observations of burned area.

We performed two model experiments to test the impact of representing the socio-economic factors on fire ignition and suppression in INFERNNO. A control experiment referred to as JULES-INFERNNO, and a similar experiment, representing the socio-economic factors through HDI on fire ignition and suppression parametrisation described in section 2.3, referred as to
350 JULES-INFERNNO+HDI.

Socio-economic impacts on fire are not represented in the initial formulation of INFERNNO described in Mangeon et al. (2016) for the ignitions and suppression of fires, which affects the $\overline{BA_{PFT}}$ values used in the initial implementation of INFERNNO. Posterior work by Andela et al. (2019) shows that average burnt area values can be larger than the ones used in the work of Mangeon et al. (2016).

355 When the HDI-based parametrisation of socio-economic impacts on fire is included in INFERNNO, it reduces ignitions and suppression of fires. Therefore, the values of $\overline{BA_{PFT}}$ are adapted to align with those reported by Andela et al. (2019). This was achieved by deriving the PFT-specific values by matching those reported by Andela et al. (2019) to the PFT categories represented in JULES as much as possible. These are not direct comparisons and a balance between the PFT representation

Table 1. Biomass burning average burnt areas ($km^2 fire^{-1}$) for all JULES plant functional types used in this configuration. Based on Burton et al. (2019) (top row) and adapted from Andela et al. (2019) (bottom row).

	Broadleaf tree			Needleleaf tree		C3			C4			Shrubs	
	Deciduous	Evergreen		Evergreen	Deciduous	Grass	Crop	Pasture	Grass	Crop	Pasture	Deciduous	Evergreen
$\overline{BA_{PFT}}$	1.7	Tropical	Temperate	1.7	1.7	3.2	0.4	3.2	3.2	0.4	3.2	2.7	2.7
Revised $\overline{BA_{PFT}}$	5.2	1.4	2.5	5.2	5.2	10.2	1.4	1.4	10.2	1.4	1.4	5.1	5.1

of the region in JULES was used to estimate a reasonable average burnt area for INFERNO. The $\overline{BA_{PFT}}$ values in both
360 experiments are detailed in Table 1.

Although INFERNO has been used in other studies to estimate fire-related emissions, this manuscript focuses exclusively on assessing the influence of the Human Development Index (HDI) on fire activity and on evaluating model performance in simulating burned area fractions. Fire emissions are therefore not analysed in this study.

3 Results

365 3.1 Representing Burnt Area - HDI relationship in INFERNO

The posterior fit distributions of the Bayesian Linear Regression parameters slope, intercept and sigma (residual variability caused by none HDI drivers) in Figure 3 and 7 show narrow intervals for the posterior parameters, evidence of high confidence in the fit for a linear relationship between HDI and deweathered burnt area fraction, with the mean Bayesian fit (dashed black line) presenting a slope of $-6.57 (\%^{-1})$, and an intercept of $19.42 (\%)$.

370 In summary, this analysis reveals a strong relationship between the HDI and burnt area (Figure 7). It demonstrates a predominantly negative correlation between HDI and deweathered burnt area globally, and especially in Eurasia, continental North America, central Asia, Southern Africa, and Australia, where increases in FWI have not translated into higher burnt areas (Figure 2). Conversely, positive correlations persist in regions like the boreal areas of North America, Siberia, and the Cerrado of South America. This evidence suggests that HDI can regionally capture part of the variability in burned area, reflecting some
375 influence of socio-economic factors.

To evaluate the use of HDI as a proxy for representing socio-economic factors influencing fire ignition and suppression in INFERNO, the methodology described in Section 2.2 was applied, and the results are presented in Figure 6, and the distribution of its posterior fits can be found in Figure 7.

380 As expected, JULES-INFERNO (the original version of the model not including the HDI) does not present a strong relationship between the deweathered burnt area fraction and HDI. The mean Bayesian fit slope is $-2.75 (\%)$, indicating a weaker relationship between the variables than in JULES-INFERNO+HDI when compared to observations. In contrast, the stronger negative slope of $-6.57 (\%^{-1})$ found in observations suggests a more pronounced socio-economic influence on fire suppression in the observational data. When HDI is explicitly incorporated to represent socio-economic effects on fire in JULES-

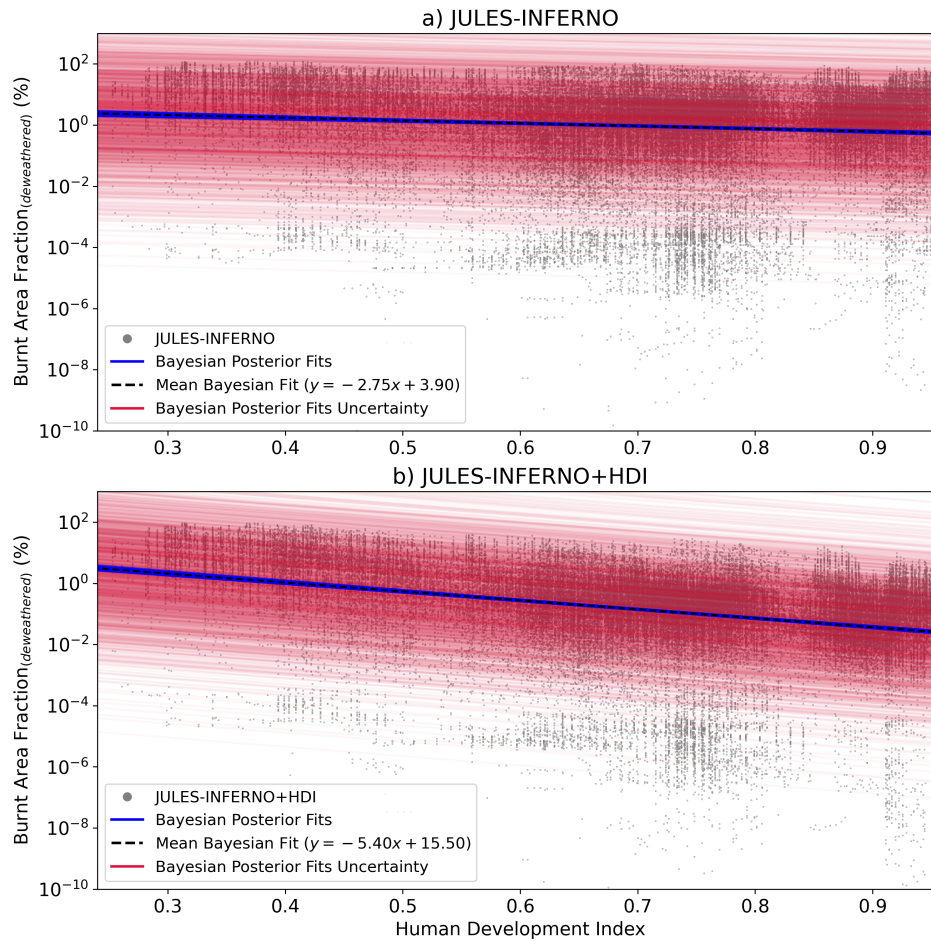


Figure 6. Relationship between the monthly mean (1997–2016) HDI and the burnt area fraction (%) for (a) JULES-INFERNO and (b) JULES-INFERNO+HDI. The grey dots represent the scatter plot for all monthly burnt area values at all grid cells as a function of HDI. The blue solid lines indicate the Bayesian regression posterior fits, while the black dashed line shows the mean Bayesian fit. The red solid lines depict the uncertainty from the Bayesian posterior fit residuals. Although the Bayesian regression was performed in log-transformed space, the results are presented here in natural space for interpretability.

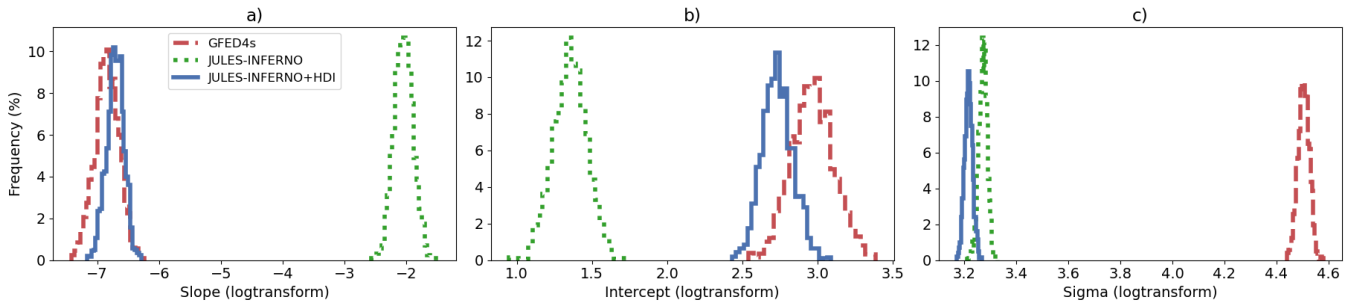


Figure 7. Histograms of posterior distributions for Bayesian linear regression fit parameters derived for GFED4s (dashed red line), JULES-INFERNO (dotted green line), and JULES-INFERNO+HDI (solid blue line). Panel a) shows the slope parameter, b) the intercept parameter, and c) depicts the sigma (error term representing random sampling noise) in log-transformed space.

INFERNO+HDI, the model better reproduces the relationship observed in GFED4s, with a Bayesian fit slope of $-5.40 (\%^{-1})$.
 385 This result demonstrates an improved representation of the socio-economic impact on fire dynamics, aligning the model more closely with observations in terms of global mean dependence on HDI.

3.2 Evaluation of impact on burnt area

To better understand the regional impact of implementing the socio-economic factors on fire ignition and suppression in INFERNO, we focus on the burnt area results averaged over the GFED4s regions as defined in Figure A1.

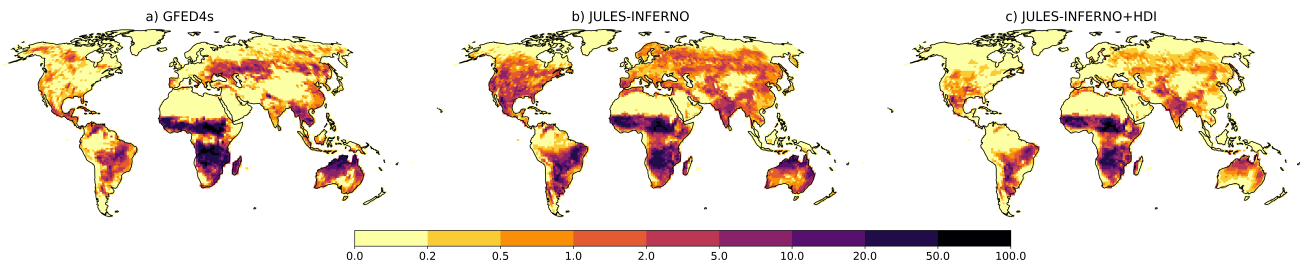


Figure 8. Burnt area fraction (%) mean annual average (1997 - 2016) for a) GFED4s, b) JULES-INFERNO and c) JULES-INFERNO+HDI. Please note that the colour mapping uses a colour axis in which the difference in colours do not correspond linearly to differences in burnt area fraction.

390 Both model experiments reproduce the overall geographical pattern of the annual average burnt area fraction (Figure 8), though with some regional differences compared to observations. For instance, JULES-INFERNO simulates the observed pattern in the major fire regions: South America, Africa and Eurasia. The 2-D cross-correlation was used to determine what is referred to as spatial correlation between the model experiments and the observation data. JULES-INFERNO shows substantial spatial biases over North America, Europe and Asia, leading to a low global spatial correlation of 0.265 compared with
 395 GFED4s. Conversely, JULES-INFERNO+HDI reduces fires in the regions with higher HDI values, reducing the biases seen in

JULES-INFERNO and resulting in a better agreement with GFED4s. JULES-INFERNO+HDI has a global spatial correlation of 0.465 when compared with GFED4s. However, due to the nature of the HDI data, sharp boundaries between countries can appear in the burnt area results (e.g., between Canada and the United States of America) - 1 e).

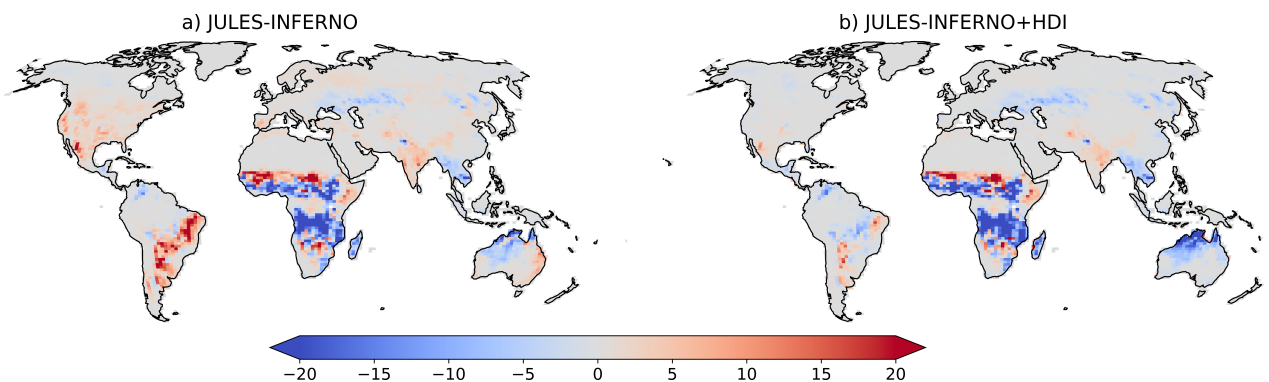


Figure 9. Burnt area fraction (%) mean annual average bias (1997 - 2016) for a) JULES-INFERNO, and b) JULES-INFERNO+HDI, calculated relative to the GFED4s observations

Representing the socio-economic factors through HDI in the parametrisation for fire ignition and suppression has an impact in all regions. However, it mostly reduces the burnt area in regions with high prosperity (high values of HDI), leading to reduced bias over North America, Europe and Asia, as shown in Figure 9. Moreover, compared to JULES-INFERNO, JULES-INFERNO+HDI reduces the positive bias over South America and India, although it increases the negative bias over the boreal regions, Australia, and South East Asia.

To further evaluate the impact of incorporating socio-economic factors on fire dynamics in INFERNO via HDI, we present histograms depicting the frequency distribution of burnt area across different fire regions in Figure 10. These histograms provide valuable insights into how the inclusion of socio-economic factors in JULES-INFERNO+HDI influences the occurrence of fires of different magnitudes compared to both JULES-INFERNO and GFED4s observations. To provide a quantitative assessment of model performance, the Wasserstein distance (Ramdas et al., 2017) is used as a metric of the fit between the probability distributions of the JULES-INFERNO and JULES-INFERNO+HDI histograms and the GFED4s observations. In this context, smaller values indicate better agreement between modelled and observed distributions, with a value of zero representing a perfect match and progressively larger values reflecting increasing divergence. This analysis helps to identify the dominant fire sizes in each region, assess whether the introduction of socio-economic factors leads to shifts in these distributions, and determine whether the implementation in JULES-INFERNO+HDI leads to a better representation of the burnt area probability distribution.

Globally, GFED4s observations display a steep decline in burnt area frequency as fire sizes increase, with small burnt area fractions dominating the distribution (Figure 10a). JULES-INFERNO reproduces this dominance of small fires relatively well but exhibits a strong underestimation of the frequency of large burnt area fractions (> 0.5). The introduction of HDI in JULES-INFERNO+HDI increases the frequency of the largest burnt area fractions (0.7–1.0), improving the representation of larger

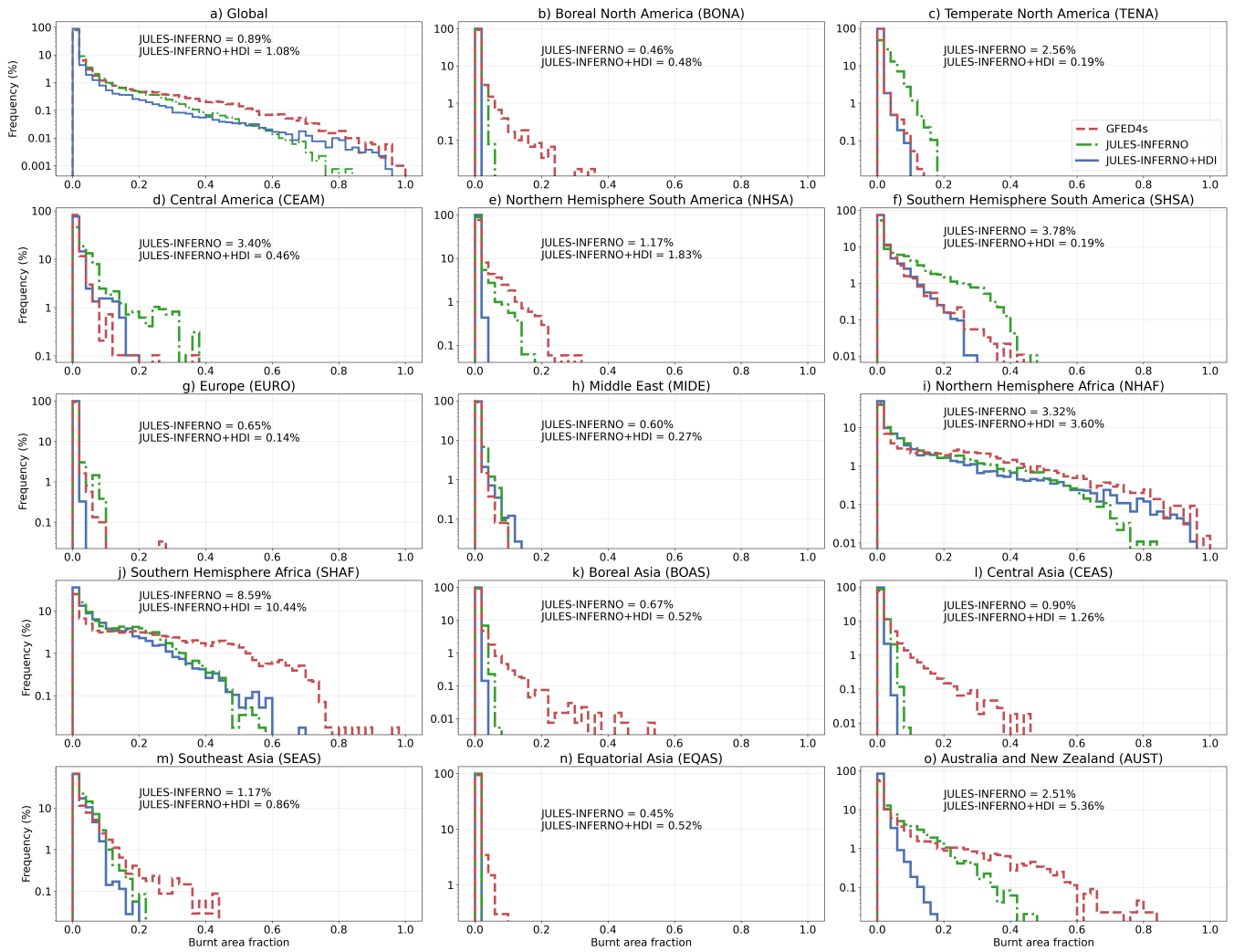


Figure 10. Histograms showing the distribution of burnt area fractions across fire regions for GFED4s observations (red dashed lines), JULES-INFERNO (green dotted lines), and JULES-INFERNO+HDI (blue solid lines), for the different fire regions. Annotated values indicate the Wasserstein distance of each model experiment relative to GFED4s.

fire sizes relative to JULES-INFERNO. However, this improvement comes at the expense of degraded performance for small
420 to moderate burnt area fractions (< 0.5), which account for the majority of fire occurrences. This trade-off is reflected in the
Wasserstein distance, which increases slightly from 0.89 % for JULES-INFERNO to 1.08 % for JULES-INFERNO+HDI at
the global scale, indicating that despite the improved fit for extreme fires, this does not compensate for the deterioration across
more frequent smaller fires.

In boreal regions (Boreal North America, BONA, and Boreal Asia, BOAS), GFED4s distributions are strongly skewed to-
425 wards small burnt areas, with a rapid decline in frequency beyond 0.1. Both JULES-INFERNO and JULES-INFERNO+HDI
substantially under-represent moderate size fires in these regions (burnt area fractions between 0.1 and 0.4), indicating per-
sistent challenges in capturing boreal fire dynamics. The inclusion of HDI leads to a modest additional suppression of burnt
areas, but the resulting Wasserstein distances remain very similar between the two model configurations (0.46 % for JULES-
INFERNO and 0.48 % for JULES-INFERNO+HDI in BONA and 0.67 % for JULES-INFERNO and 0.52 % for JULES-
430 INFERNO+HDI in BOAS). This indicates that incorporating socio-economic factors does not provide a clear improvement
or degradation in the overall agreement with observations in boreal regions, suggesting that other processes, such as fuel
availability, fire weather, or ignition sources, are likely more influential in governing boreal fire regimes.

In temperate regions, including Temperate North America (TENA) and Europe (EURO), GFED4s observations again show
a strong dominance of small fires (burnt area fractions < 0.1). In TENA, JULES-INFERNO markedly overestimates burnt area
435 fractions greater than 0.1, while JULES-INFERNO+HDI substantially reduces this bias. This improvement is quantitatively
supported by a pronounced reduction in Wasserstein distance from 2.56 % to 0.19 %, indicating a clear overall improve-
ment across the full distribution. For EURO, although JULES-INFERNO+HDI reduces the occurrence of larger fires, it also
underestimates small and moderate burnt area fractions. Consequently, the Wasserstein distance decreases from 0.65 % in
JULES-INFERNO to 0.14 % in JULES-INFERNO+HDI, suggesting a closer overall distributional match, but Figure 10g
440 shows that this improvement does not consistently reflect better performance across all fire size classes. In particular, JULES-
INFERNO better captures the frequency of small fires, which dominate the observed distribution, indicating that the apparent
improvement in the Wasserstein distance should be interpreted cautiously.

In Central America (CEAM) and Southern Hemisphere South America (SHSA), GFED4s exhibits broader distributions with
contributions from small to moderate burnt areas. JULES-INFERNO strongly overestimates moderate and large fires in both
445 regions, leading to large Wasserstein distances (3.40 % and 3.78 %, respectively). The inclusion of HDI substantially narrows
the distributions and reduces the occurrence of large fires, resulting in marked reductions in Wasserstein distance (to 0.46
% in CEAM and 0.19 % in SHSA). These results indicate a genuine improvement across most of the distribution, including
the dominant fire sizes. In African savanna regions (Northern Hemisphere Africa, NHAf, and Southern Hemisphere Africa,
SHAF), GFED4s shows relatively high frequencies of large burnt areas. Both experiments underestimate extreme fire sizes, but
450 JULES-INFERNO+HDI increases the suppression of medium fires further, leading to mixed results. In NHAf, the Wasserstein
distance increases from 3.32 % in JULES-INFERNO to 3.60 % in JULES-INFERNO+HDI, while in SHAF it increases from
8.59 % to 10.44 %, indicating a degradation in overall distributional agreement despite improvements in larger fire sizes.

Across Asian regions (EQAS, CEAS, SEAS) and Australia and New Zealand (AUST), GFED4s distributions are dominated by small to moderate burnt area fractions. Both experiments under-predict medium-sized burnt area fractions (e.g., 0.2–0.6),
455 with JULES-INFERNO+HDI generally exacerbates this negative bias. This is reflected in increased Wasserstein distances for JULES-INFERNO+HDI in CEAS (1.26 % for JULES-INFERNO and 0.90 % for JULES-INFERNO+HDI) and AUST (2.51 % for JULES-INFERNO and 5.36 % for JULES-INFERNO+HDI), indicating that the additional suppression associated with HDI reduces agreement with observations across the most frequent fire sizes.

Overall, the introduction of a socio-economic representation of fire suppression in JULES-INFERNO+HDI reduces the frequency of large burnt areas and corrects strong positive biases present in JULES-INFERNO in several regions, notably TENA, CEAM, and SHSA. However, these improvements are often accompanied by a deterioration in the representation of small and moderate burnt area fractions, which dominate fire occurrence globally. As reflected by the Wasserstein distance, JULES-INFERNO+HDI does not consistently outperform JULES-INFERNO across all regions, particularly in savanna-dominated regions such as Africa, Australia, and parts of Eurasia. These results highlight the need for further refinement of the socio-economic parametrisation and the inclusion of additional region-specific processes to improve the simulation of fire size distributions across all fire regimes.
465

Figure 11 shows the burnt area annual mean time series. To assess the ability of the model to reproduce the observed time series, we perform a statistical analysis based on monthly averaged data, including the calculation of various metrics such as the Root Mean Squared Error (RMSE), Root Mean Squared Error after removal of a constant mean bias ($RMSE_{UB}$), the bias, Pearson correlation, and the Standard Deviation (STD) of the burnt area monthly and annual mean time series for all GFED4s regions (Table A1). The STD is computed after removal of a linear trend. We also assess the model's ability to reproduce observed trends using a simple log-transformed linear regression. Figure 12 summarises these statistical measures for both model configurations.
470

The results presented in Figures 11 and 12, together with the regional statistics in Table A1, show that the inclusion of socio-economic factors in INFERNO leads to regional improvements in the simulation of burnt area in regions where JULES-INFERNO bias exhibits the largest deviations from GFED4s (greater than 150 %). For example, the relative bias in TENA is reduced from 735.6 % in JULES-INFERNO to 44.5 % in JULES-INFERNO+HDI, in CEAM from 259.2 % to 24.2 %, in SHSA from 191.7 % to -1.7 %, in EURO from 258.8 % to -48.8 %, and in MIDE from 420.5 % to 231.8 % (Table A1). These reductions are also accompanied by decreases in RMSE, indicating an improved agreement between JULES-INFERNO+HDI and GFED4s for these regions.
480

Conversely, some regions experience a smaller, but still notable, increase in relative bias. For example, relative bias in Northern Hemisphere South America (NHSA) changes from -60.1 % to -94.3 %, in Australia (AUST) from -21.8 % to -82.3 %, and in Southern Hemisphere Africa (SHAF) from -45.3 % to -55.3 %. This highlights that the inclusion of HDI does not uniformly improve model performance in all regions, reflecting a trade-off between correcting large positive biases and potentially amplifying negative biases in other areas.
485

In addition, the inclusion of HDI is associated with a systematic reduction in the simulated variability. The ratio of modeled to observed standard deviation (STD/STD_{GFED4s}), calculated from detrended monthly mean values, decreases in most regions.

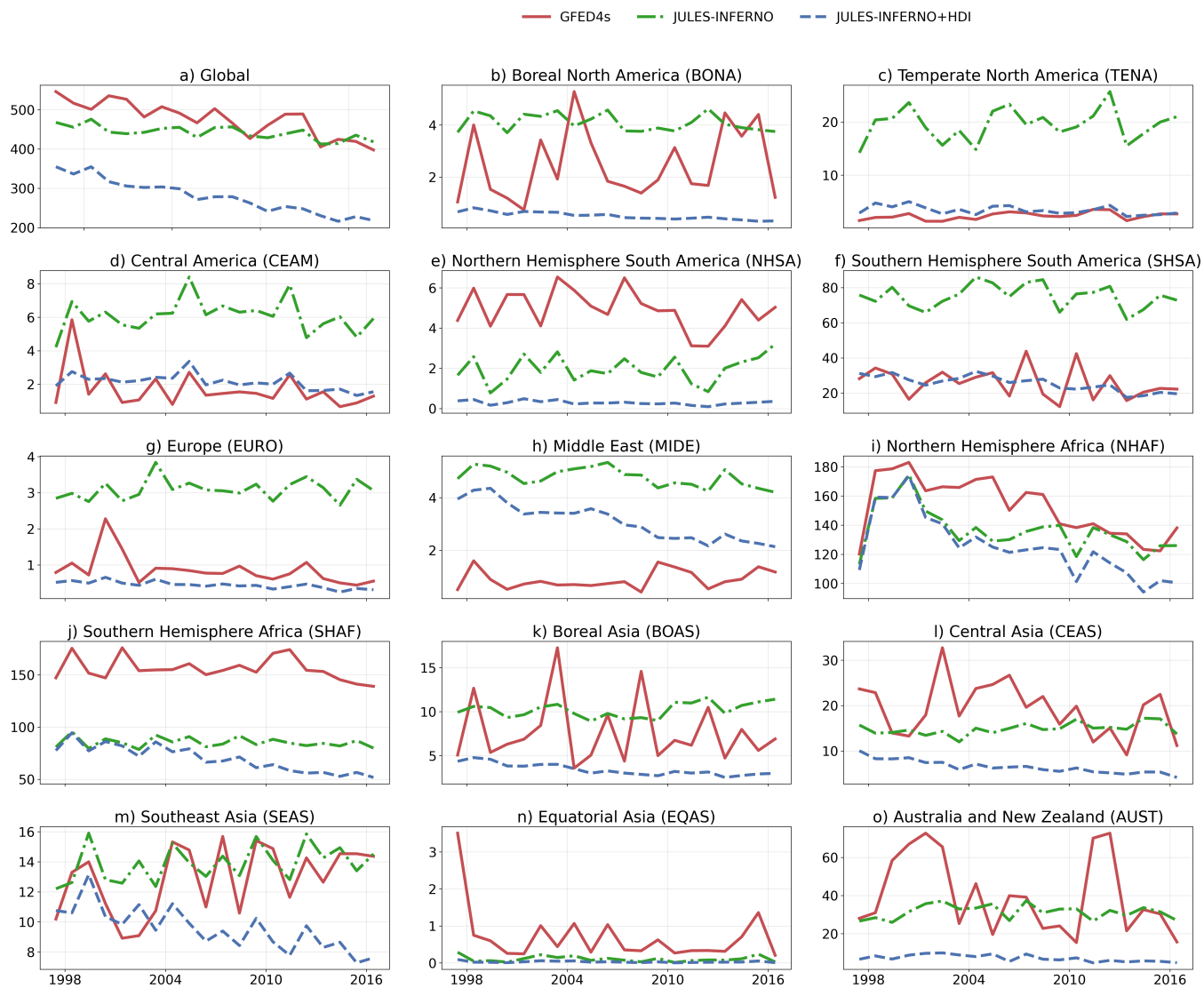


Figure 11. Time series of annual mean burned area (*Mha*) from 1997 to 2016 across different fire regions, shown for GFED4s (solid red line), JULES-INFERNO (green dot-dash line), and JULES-INFERNO+HDI (blue dashed line).

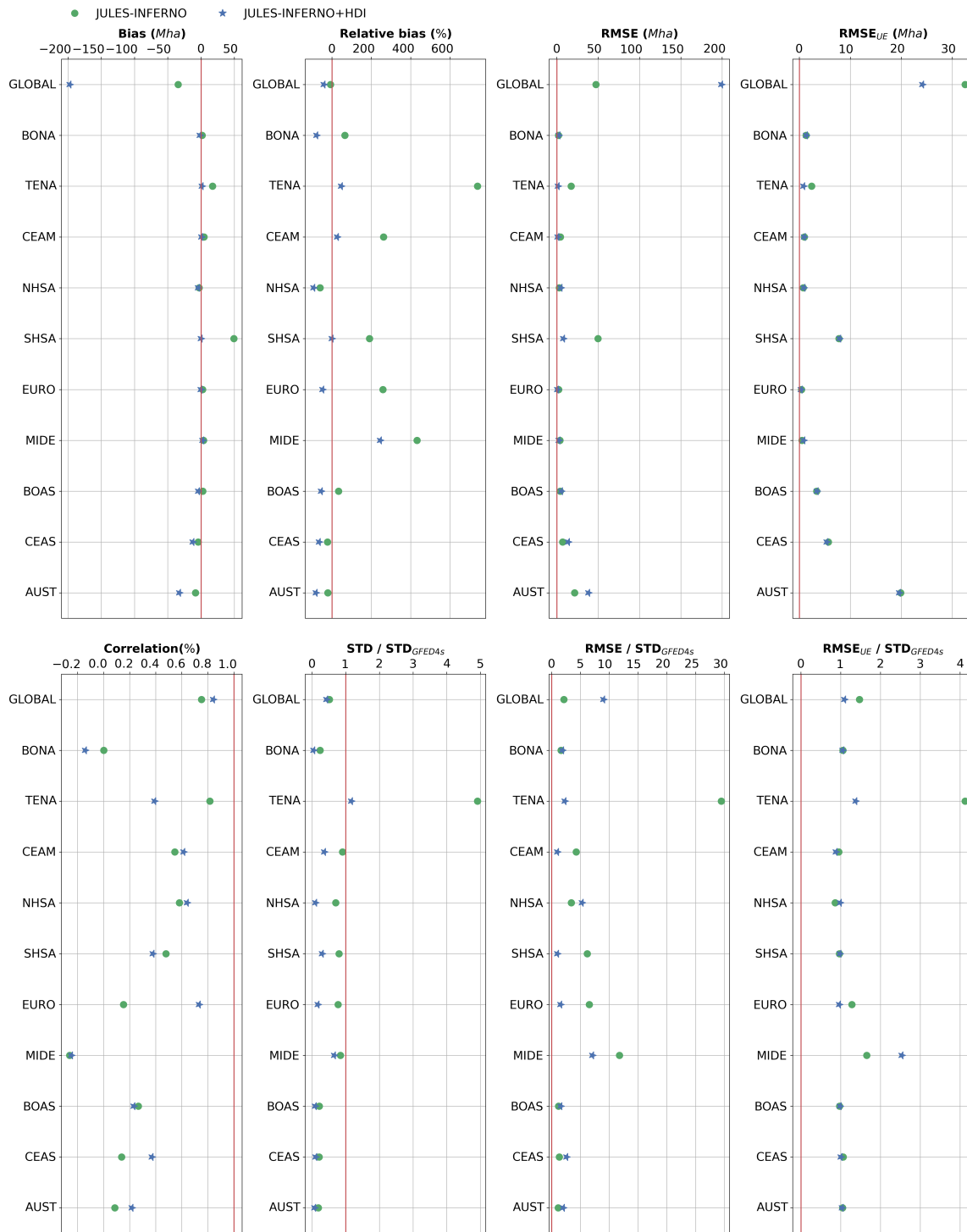


Figure 12. Summary of the statistics presented in Table A1 comparing JULES-INFERNO (green circle) and JULES-INFERNO+HDI (blue star). The red line show the reference value for a perfect simulation, the closer the experiment symbol is from this line the better.

For instance, STD/STD_{GFED4s} decreases from 4.41 to 1.04 in TENA, from 0.25 to 0.11 in BONA, from 0.84 to 0.34 in CEAM, from 0.77 to 0.28 in SHSA, and from 0.68 to 0.15 in SHSA, indicating an underestimation of sub-annual variability. While
490 this reduction in variability can contribute to lower RMSE in some regions, it also suggests that in JULES-INFERNO+HDI the amplitude of burnt area fluctuations is damped relative to observations.

At the global scale, JULES-INFERNO exhibits a relatively small mean bias due to compensating regional errors. The inclusion of HDI reduces these compensating effects, resulting in a larger negative global relative bias (from -7.21 % in JULES-INFERNO to -41.46 % in JULES-INFERNO+HDI) and an increase in RMSE (from 42.28 to 198.99). Nevertheless,
495 $RMSE_{UB}$ decreases from 32.50 to 24.12, and the Pearson correlation with GFED4s increases from 0.75 to 0.84, reflecting an improved representation of temporal co-variability of monthly burnt area.

Overall, the inclusion of socio-economic factors in INFERNO through HDI results in a trade-off. JULES-INFERNO+HDI improves the mean-state representation in regions with the largest bias in JULES-INFERNO (TENA, CEAM, SHSA, EURO, MIDE). However, relative biases worsen in regions with initial where burnt area is underestimated in JULES-INFERNO
500 (NHSA, SHAF, AUST), and variability is generally reduced.

Nonetheless, the improvements from JULES-INFERNO+HDI in regions such as TENA, NHAF, and SHAF have a greater impact on the global metrics than the reduced performance seen for regions such as CEAM, NHSA, SHSA, EURO, and MIDE. For regions such as BOAS, CEADS, SEAS, EQAS, and AUST, both model configurations underperform in terms of standard deviation, and any differences between the STD / STD_{GFED4s} are small when compared to the observed standard deviation
505 (e.g., difference between the JULES-INFERNO and JULES-INFERNO+HDI STD / STD_{GFED4s} smaller than 15 %).

Furthermore, for some of these regions INFERNO is not expected to agree well with observations, especially in terms of variability, as the fire behaviour of some of these regions is characterised by mechanisms that are not represented in INFERNO. This will be further discussed in Section 4.

3.3 Impact on burnt area trends

510 Representing the socio-economic factors through HDI in INFERNO adds a new external constraint to the model. Through this, historical changes to socio-economic factors influence how changes in population density affect fire ignitions in the model (see Section 2.3 and Figure 4). Specifically, for regions with high HDI, variations in population have less of an impact on anthropogenic ignitions, while for regions with low HDI, variations in population can have a more considerable impact. This alters the importance of population density changes for highly developed regions, making HDI the dominant factor shaping
515 burnt area trends.

As shown in Figure 13, both JULES-INFERNO and JULES-INFERNO+HDI represent the main global burnt area trends. JULES-INFERNO is able to represent the regions with burnt area increases (e.g., Southern Africa and Northeast South America) and captures the dominant region for decreased burnt area - North Africa. However, this model setup tends to have weaker negative trends when compared to GFED4s. Conversely, JULES-INFERNO+HDI presents stronger trends, better representing those found in observations. However, it does not reproduce the positive trends in Southern Africa and Northeast South
520

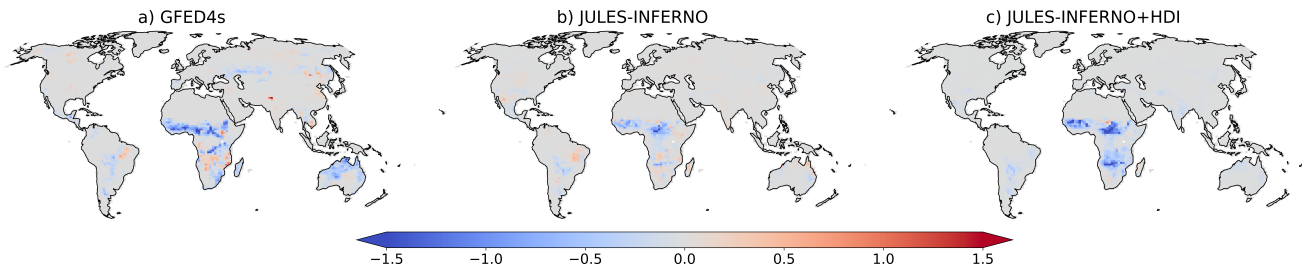


Figure 13. Burnt area fraction trend ($\% \text{ year}^{-1}$) (calculated between the period 1997 - 2016) for a) GFED4s, b) JULES-INFERNO and c) JULES-INFERNO+HDI.

America. Both JULES-INFERNO and JULES-INFERNO+HDI are unable to represent the observed trends in Central Asia or Boreal North America.

Nonetheless, it should be noted that over the 2001–2012 period, Andela and Van Der Werf (2014) estimated that 51 % of the upward trend over southern Africa can be attributed to El Niño-Southern Oscillation (ENSO), while there is also evidence that socio-economic developments can be responsible for a decline. The relation between ENSO and annual burned area depends both on the effect of ENSO on precipitation and on the antecedent precipitation-burned area response. While the model setup is able to capture ENSO variability, as its weather is driven by reanalysis, there is no mechanism that allows INFERNO to represent the antecedent precipitation-burned area effects due to litter build up. This is a limitation of the model and it should not be expected for the model to perform well in regions where this precipitation-burned area coupling can be dominant, such as Central America, Northern Hemisphere South America, Europe, Northern Hemisphere Africa, and Central Asia (Andela et al., 2017; Abatzoglou et al., 2018).

Overall, representing the socio-economic factors through HDI in INFERNO results in an improvement in burnt area trends in comparison with observations. As seen in Table A1, JULES-INFERNO+HDI better represents the global negative trend in burnt area when compared to observations ($-6.77 \text{ Mha year}^{-1}$ for GFED4s, $-2.24 \text{ Mha year}^{-1}$ for JULES-INFERNO, and $-7.58 \text{ Mha year}^{-1}$ for JULES-INFERNO+HDI). This improvement comes mostly from a better representation of the burnt area trends in regions with strong negative trends, such as SHSA, NHAF, CEAS and AUST, but also by better representing regions with weak negative burnt area trends, namely CEAM, NHSA, EURO and BOAS. Moreover, in regions such as CEAM, NHSA, EURO, BOAS, CEAS, and AUST, JULES-INFERNO+HDI shows a negative burnt area trend, in better agreement with observations.

Contrary to these improvements, JULES-INFERNO+HDI can also produce trends that are too strong, reflecting the presence of compensating biases in the model. For example, in Southern Hemisphere Africa (SHAF), although JULES-INFERNO+HDI reproduces the observed negative burnt area trend, the magnitude of the trend is substantially overestimated ($-0.54 \text{ Mha year}^{-1}$ for GFED4s, $-0.14 \text{ Mha year}^{-1}$ for JULES-INFERNO, and $-1.94 \text{ Mha year}^{-1}$ for JULES-INFERNO+HDI). In this region, the improved agreement in the sign of the trend in JULES-INFERNO+HDI arises from compensating biases that mask an excessive sensitivity to human influence, whereas JULES-INFERNO provides a more realistic representation of the trend magnitude. This

example highlights that, as for JULES-INFERNO, apparent improvements in some metrics in JULES-INFERNO+HDI, such as regional or global trends, can result from compensating errors rather than an overall improvement in process representation.

Nonetheless, the observed dataset (GFED4s) shows that out of 14 regions, four have positive burnt area trends (Table A1). JULES-INFERNO only presents a positive trend for TENA and SEAS. While JULES-INFERNO+HDI tends to enforce decreasing trends, this only happens in four regions out of 14 (i.e., TENA, SHAF, MIDE, and SEAS). For the remaining 10 regions, JULES-INFERNO+HDI presents a similar trend to JULES-INFERNO or even an improved trend when compared to GFED4s.

It should be noted that, in some of these regions INFERNO does not model all the processes that represent fire behaviour. This has an impact on overall model results. For example, due to the typical model resolution and timescales in Earth System Modelling, INFERNO was not designed to model the processes and mechanisms that are needed to represent large and severe fires which dominate the trends and fire regime characteristics of these regions. Therefore, it is expected that regions where fire regimes are dominated by large and severe fires may be affected by a negative bias in burnt areas and fire emissions, as well as on their response to a changing climate.

3.3.1 Impact of external model drivers on burnt area trends

As described in section 2.4, the JULES-ES experimental setup relies on ancillary forcing data to represent external processes to JULES, such as atmospheric weather conditions, atmospheric composition, population density, and biogenic drivers. These external forcings can drive fire by forcing changes to the evolution of land surface properties, fire ignitions, and fire weather. It is important to understand the impact these external drivers have on the burnt area trends and the interaction with the parametrised socio-economic factors in fires. Therefore a set of sensitivity experiments was performed by fixing the external model drivers to the year 1990 and only allowing an individual external driver to vary transiently through the experiment.

- **1990 control:** where all external model drivers are fixed to year-1990 values
- **clim:** where only the atmospheric drivers are transient (downward longwave radiative flux, downward shortwave radiative flux, precipitation, surface pressure, air temperature, meridional and zonal wind components)
- **tas:** where only air temperature at 2 m is transient
- **ppn:** where only precipitation is transient
- **lu:** where only the land use is transient
- **Ndep:** where only Nitrogen deposition is transient
- **pop:** where only population density is transient
- **CO₂:** where only the atmospheric carbon dioxide (CO₂) mixing ratio is transient
- **HDI:** where only the Human Development Index is transient (only for JULES-INFERNO+HDI)

These sensitivity experiments branched from their respective control runs - JULES-INFERNO and JULES-INFERNO+HDI - starting from 1990 and run up to 2016. In this way, the underlying land surface state from the reference run is preserved, and only changes to the forcing that take place during the period of interest are taken into effect. The trends for each relevant external forcing are in Supplementary Figure A2.

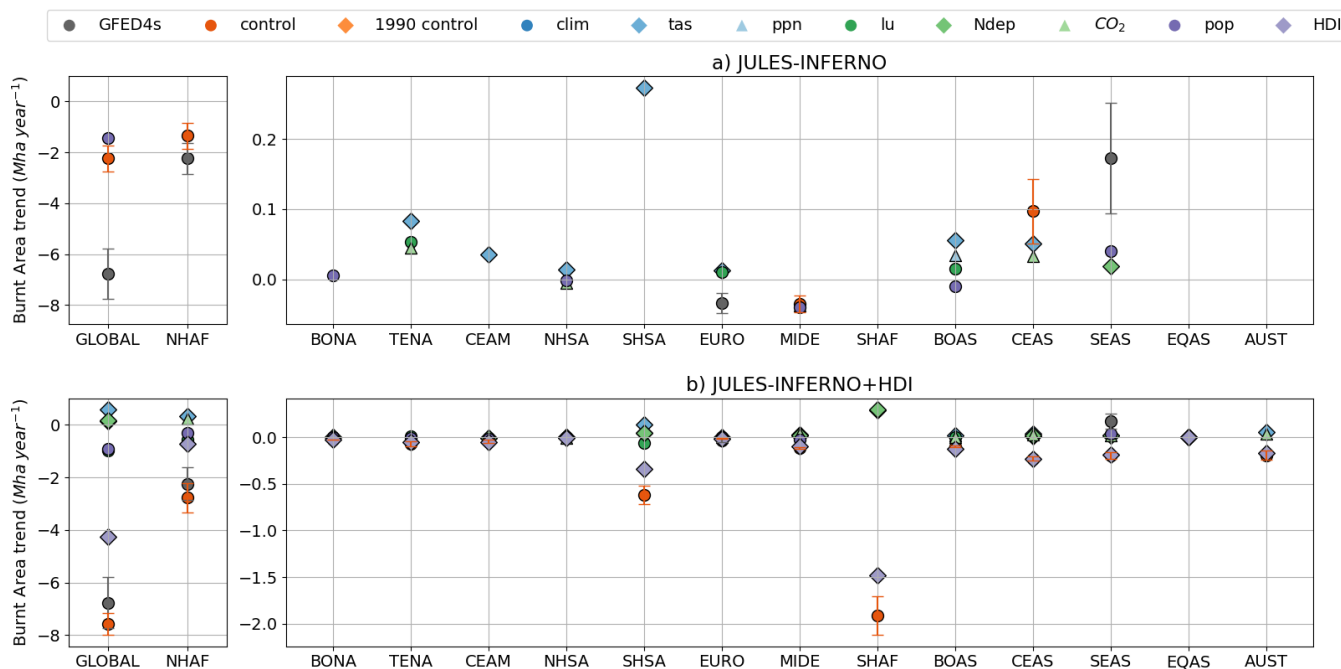


Figure 14. Burnt area trends ($Mha; year^{-1}$) for the different GFED4s fire regions (Giglio et al., 2013) from the model sensitivity experiments for a) JULES-INFERNO and b) JULES-INFERNO+HDI. Only trends that are statistically significant at the 95 % confidence level are shown. The whiskers for the values of GFED4s and *control* represent the standard error associated with the trend. GLOBAL and NHAF regions are presented in separate panels to improve readability, due to their large trends compared to other regions.

580 The results of these sensitivity experiments on the burnt area trends ($Mha; year^{-1}$), and respective standard error, for the different GFED4s fire regions are presented in Table A2 in the Appendix for JULES-INFERNO and Table A3 for JULES-INFERNO+HDI. These results are summarised in Figure 14. Burnt area trends in JULES-INFERNO tend to be driven by climate, land use, or population density changes (relative contribution greater than 50 % when compared to their reference), with the dominant driver (the sensitivity experiment with the largest absolute trend value) for the majority of regions being climate (including through air temperature and precipitation), for example, BONA, TENA, NHSA, SHSA, EURO, MIDE, SHAF, BOAS, CEAS, SEAS, and AUST. On the one hand, air temperature is a dominant, and statistically significant, driver of increasing burnt area trends for CEAM, NHSA, BOAS and CEAS. For MIDE, precipitation has a dominant role in reducing burnt area. On the other hand, despite dominating the burnt area trends, temperature can have opposite effects. Namely, for 585 TENA, SHSA, EURO, and AUST, where temperature results in an increase in the trend.

590 Anthropogenic drivers, such as land use and population density can play a major role in some regions. For example, land use is the dominant driver in SHSA and SHAF, population density in CEAM, and both are important drivers in TENA, EURO, MIDE, NHAF. Land use can cause either an increase (TENA, EURO, and SHAF) or a decrease in burnt area, while population density results in a reduction in burnt area for all the regions where this external driver is dominant (TENA, CEAM, EURO, MIDE, and NHAF).

595 Biogenic drivers such as nitrogen deposition and atmospheric carbon dioxide assimilation tend to play a less significant role in the burnt area trends, impacting burnt area trends only for NHSA through carbon dioxide, and CEAS through both nitrogen deposition and atmospheric carbon dioxide.

For JULES-INFERNO+HDI, HDI is the dominant factor in the burnt area trend for all regions, with the exception of NHAF and EQAS. This is evident in the sensitivity experiments for JULES-INFERNO+HDI where only HDI forcing is made transient
600 (*HDI*). In all cases, representing the socio-economic factors through HDI in INFERNO changes the relative role that external forcings have in determining burnt area trends. For example, it reduces the role that climate, population density, and land use have in burnt area trends towards a stronger role from socio-economic effects (HDI).

This impact is especially evident when comparing the role of the climate forcing in driving burnt area trends between JULES-INFERNO and JULES-INFERNO+HDI in their respective sensitivity experiments, *clim*, *tas*, and *ppn*. For instance,
605 for regions where temperature effects on burnt area trends were dominant in JULES-INFERNO, they show less of an impact from temperature in JULES-INFERNO+HDI, albeit still at a statistically significant level (e.g., TENA, EURO, CEAS, and AUS). In addition, where the climate contributions to the burnt area trends were small in JULES-INFERNO (BONA, CEAM, and NHSA), when considering socio-economic factors, these become statistically non-significant in terms of the climate contributions to the trends. In addition, representing the socio-economic factors through HDI in INFERNO leads to a change
610 in how precipitation and temperature can impact some regions. For example, precipitation was the dominant driver of burnt area trends in JULES-INFERNO for MIDE, causing a reduction of burnt area (trend of $-0.038 \text{ Mha year}^{-1}$). Conversely, in JULES-INFERNO+HDI, precipitation has less impact on the burnt area trend (trend of $-0.016 \text{ Mha year}^{-1}$), while temperature has a larger impact on the burnt area trend ($+0.034 \text{ Mha year}^{-1}$). A similar result is also seen for NHAF, where the role of temperature becomes statistically significant in JULES-INFERNO+HDI (trend of $+0.342 \text{ Mha year}^{-1}$).

615 Moreover, results show a difference in the role of anthropogenic drivers (land use and population density) on burnt area trends between JULES-INFERNO and JULES-INFERNO+HDI sensitivity experiments. For BONA, CEAM, and NHSA regions, burnt area decreases for JULES-INFERNO+HDI (*pop* and *lu*) sensitivity experiments while for JULES-INFERNO (*pop* and *lu*) an increase was simulated.

Together, these results indicate that representing socio-economic factors through HDI can influence the simulation of burnt
620 area in ESMs, but that its impact is regionally heterogeneous and involves clear trade-offs. The inclusion of HDI in JULES-INFERNO primarily improves the mean-state representation and trend magnitude in regions where the original model exhibits strong positive biases. However, these improvements are not uniform across all fire regions. In several regions the additional constraint imposed by the use of HDI as a proxy for socio-economic impacts on fire leads to reduced variability or amplified negative biases. Overall, JULES-INFERNO+HDI demonstrates that socio-economic development can be an important modu-

625 lator of fire activity and trends at regional scales, but further refinement of JULES-INFERNO+HDI is still required to better
balance its effects across different fire regions.

4 Discussion & Conclusions

This work aims to represent socio-economic factors, through the use of the HDI, together with the Pechony and Shindell (2009)
anthropogenic fire ignitions to parameterise human socio-economic impacts on fires. When using the INFERNO fire model,
630 the aim was to improve the regional representation of human–environmental coupling for applications at large spatial scales
within an ESM.

The results presented in this study indicate that incorporating socio-economic factors through HDI in the fire ignition and
suppression parametrisation within INFERNO, alongside the revised $\overline{BA_{PFT}}$ parameters, can improve the representation of
the globally-averaged linear relationship between burnt area and HDI, as observed in Figures 3 and 6. This adjustment leads to
635 better performance in regions where JULES-INFERNO previously exhibited large positive biases. However, the improvements
are regionally variable, and in some areas the inclusion of HDI introduces trade-offs, such as reduced variability or less accurate
representation of small to moderate fires, highlighting the need for further refinement of the approach.

The simple representation of HDI captures how more complex socio-economic processes tend to suppress fires as develop-
ment increases. In our analysis, we modelled the relationship between burnt area and HDI using a linear form for simplicity. It
640 is important to note that this does not imply that the relationship is inherently linear in the observed data. This result should be
interpreted as conditional on the linear model; it reflects a statistical trend rather than a directly observed linear relationship.

This approach suggests that regions with higher HDI tend to have lower average values of burnt area. This trend can be
attributed to various components of HDI that influence government policies and resources for fire management (Miranda-
Lescano et al., 2023). For instance, the income component of HDI broadly reflects economic capacity, which can affect the
645 ability to resource wildfire management; accordingly, Rideout et al. (2017) illustrate how strategic budgeting and planning
tools (e.g., STARFire) can be used to allocate limited resources across preparedness and fuel treatment programmes based
on risk, valuation, and cost information. Rizzo and Rizzo (2024) highlight that wildfire smoke has substantial adverse health
impacts and therefore motivates mitigation and adaptation actions (e.g., risk communication and preparedness/early-warning
measures) aimed at reducing exposure and protecting vulnerable populations. Additionally, the education component of HDI
650 may be relevant because public awareness and preparedness can influence fire risk; for example, a Florida case study found that
wildfire prevention education efforts (e.g., public service announcements and outreach) were associated with fewer preventable
human-caused ignitions across several ignition categories (Prestemon et al., 2010).

Collectively, these factors help explain the observed trend and underscore the utility of HDI as a general proxy for represent-
ing socio-economic influences in fire management, but may not be fully representative of specific particularities of any single
655 given region.

Large bias reductions are evident in Temperate North America (TENA), Central America (CEAM), Southern Hemisphere
South America (SHSA), Europe (EURO), and Middle East (MIDE), with the largest reductions in TENA where a 735.57

% bias in JULES-INFERNO is reduced to 44.46 % in JULES-INFERNO+HDI. It should be noted that for Australia and New Zealand (AUST) and East Asia (SEAS), JULES-INFERNO+HDI performance is degraded, increasing the negative bias compared to JULES-INFERNO. JULES-INFERNO includes contributions from regions with large positive biases exceeding +150 % (TENA, CEAM, SHSA, EURO, and MIDE), with a combined bias of +76.96 Mha (Table A1). These are the regions targeted for bias reduction in JULES-INFERNO+HDI. Considering that JULES-INFERNO's overall global bias is -34.35 Mha, removing the compensating effect of these highly biased regions would imply a potential global bias of approximately -111.31 Mha. This highlights the importance of addressing regional biases and demonstrates that although JULES-INFERNO appears to perform well at the global scale—as seen when comparing annual mean burnt area against GFED4s in Figure 12, this is largely due to compensating errors at the regional level.

The histograms of burnt area frequency across different fire regions (Figure 10) reveal key differences in how fire sizes are distributed between JULES-INFERNO, JULES-INFERNO+HDI, and GFED4s observations. The implementation of HDI leads to improvements in regions such as Temperate North America (TENA), Central America (CEAM), and Southern Hemisphere South America (SHSA), where reductions in the frequency of large fires correct strong overestimations present in JULES-INFERNO. In other regions, such as Northern Hemisphere Africa (NHAF), the inclusion of HDI further suppresses medium to large fires, but this does not consistently improve agreement with observations. In Northern Hemisphere South America (NHSA) and Australia and New Zealand (AUST), the underprediction of medium and large fire sizes in JULES-INFERNO further degrades with the inclusion of HDI, highlighting the regional trade-offs associated with the socio-economic parametrisation.

Moreover, representing the socio-economic factors through HDI in INFERNO-JULES+HDI improves the representation of the burnt area trends, especially in areas where GFED4s presents negative trends. At the same time, JULES-INFERNO shows no significant trends (e.g., SHSA, NHAF, CEAS, and AUST), as well as better representing regions with weak negative burnt area trends (CEAM, NHSA, EURO, and BOAS). However, JULES-INFERNO+HDI can also produce overly strong trends (e.g., SHAF) or misrepresents the observed positive burnt area trends found in TENA, MIDE, BONA, and SEAS.

As mentioned previously, observations (GFED4s) show that out of 14 regions, four have a positive burnt area trends. From these, only JULES-INFERNO shows a positive trend for TENA and SEAS. While JULES-INFERNO+HDI tends to strengthen decreasing trends, this only happens in four regions out of 14 (TENA, SHAF, MIDE, and SEAS). For the remaining 10 regions, JULES-INFERNO+HDI shows a similar trend to JULES-INFERNO or even an improved trend when compared to GFED4s.

It should be highlighted that in some of these regions, INFERNO does not model all of the processes that impact fire behaviour. This has an impact on overall model results. For example, INFERNO was not designed to capture the dynamics of large, severe fires that dominate fire regimes in some regions. As a result, these areas may show a negative bias in burned area and fire emissions, as well as in their response to climate change.

Overall, the improved representation of the burnt area trends in JULES-INFERNO+HDI when compared to JULES-INFERNO highlights the importance of including the socio-economic factors in fire ignition and suppression in order to better reproduce the observed fire trends. This impact is especially evident when comparing the role of external climate drivers on burnt area trends between JULES-INFERNO and JULES-INFERNO+HDI using a set of sensitivity experiments. The results of these

experiments show that including socio-economic impacts on fire results in the burnt areas trends being dominated by socio-economic drivers through a reduction in the contribution from climate drivers, especially from temperature and precipitation.

695 **4.1 Modelled burnt area trends**

We have shown that introducing the representation of socio-economic factors can change the impact external forcing has on burnt area trends and that the mechanisms that lead to this can differ at a regional level. For example, the inclusion of socio-economic factors reduces the role of temperature in driving trends (e.g., increase for TENA, EURO, CEAS, and AUS), as well as by changing the behaviour that climate drivers have in burnt area trends (e.g., MIDE, NHAF, and SEAS). Socio-economic factors also alter the influence of land use and population density, increasing their impact on fire activity by shaping how human activity interacts with fire regimes. (e.g., BONA, CEAM, and NHSA).

Although HDI does not explicitly encompass the impacts of fire management policies, these results are consistent with previous studies demonstrating that, in highly developed regions, institutional capacity and land-management policies exert a stronger control on fire activity than individual ignition behaviours. For example, Carreiras et al. (2014) show that in Mediterranean Europe, changes in land management and fire suppression policies have contributed to declining burnt area despite sustained human presence. Mourão and Martinho (2014) highlight the role of coordinated public policies and investment in fire prevention and suppression infrastructure in reducing fire impacts in developed regions. Ford et al. (2021) demonstrate that governance quality and fire management strategies, rather than ignition pressure alone, are key determinants of fire outcomes in high-income countries. Similarly, Jacobson et al. (2022) find that institutional fire management capacity strongly moderates the relationship between human activity and fire occurrence. Finally, Nikolakis and Roberts (2022) emphasise the importance of governance structures and policy frameworks in shaping fire regimes, showing that effective land and fire management institutions can substantially limit fire extent even under increasing anthropogenic pressure.

The work of Kelley et al. (2019) and Jones et al. (2022) shows that, despite the increases in fire weather seasons and fire weather extremes that have been observed in all world regions, burnt area has shown a variety of regional trends and that the negative trends were found to be significant only in Africa (NHAF and SHAF), Europe (EURO), and Central Asia (CEAS). At a global scale, burnt area trends show a decline predominantly driven by a decline in burnt area in the savannah-grassland systems caused by the expansion of high-capital agriculture (Andela et al., 2017), as well as reductions in vegetation productivity driven by changes to the hydrological balance (Zubkova et al., 2019). The results on the impact of external model drivers on burnt area trends, detailed in Section 3.3.1, agree with this. In both JULES-INFERNO and JULES-INFERNO+HDI, the dominant factor contributing to the negative trend in global burnt area are linked to anthropogenic drivers (land use, population density and HDI), as well as precipitation.

Several studies have shown that declines in burnt area across the Mediterranean have occurred despite increases in fire weather severity and extensions of the fire weather season, with these declines largely attributed to enhanced fire prevention, suppression capacity, and fire management strategies. For instance, (Rabin et al., 2017) demonstrate that improvements in suppression efficiency and fire management practices have offset the effects of increasingly severe fire weather in Mediterranean Europe. (Jones et al., 2022) show that trends in burnt area are weakly coupled to fire weather in recent decades, highlighting

the dominant role of human intervention in limiting fire spread. Similarly, (Carreiras et al., 2014) attribute declining burnt area to changes in land management and fire suppression policies rather than to climatic drivers alone. In addition, (Mourão and Martinho, 2014) emphasise that increased investment in fire prevention and coordinated policy responses have been critical in
730 mitigating fire impacts despite a lengthening fire season.

Moreover, results show the impact of the anthropogenic drivers (land use and population density) have in the burnt area trends, resulting in a decrease in burnt area for JULES-INFERNO+HDI compared to JULES-INFERNO, for BONA, CEAM, and NHSA regions. This result, combined with the impact seen in the reduction of the effects temperature has in burnt area, are especially relevant in South America (NHSA and SHSA), leading to a better performance of JULES-INFERNO+HDI in
735 these regions overall. The impact of socio-economic effects on fire is well documented for the Amazonia region, where fire is routinely used for land clearing and is tightly linked to deforestation rates (Silva Junior et al., 2021). Over the last decade, declines in deforestation rates owing to policy interventions and regulatory enforcement have led to an observed negative trend in overall burnt area (Nepstad et al., 2014). However, this decline in burnt area has not been uniform, as drought-related fire incidence has sometimes increased even when deforestation slowed, reflecting historical shifts in economic conditions and
740 environmental policies (Aragão et al., 2018).

The analysis of the impact of representing the socio-economic factors through HDI in INFERNO has also shown that JULES-INFERNO+HDI can also producing overly strong negative trends resulting in worse performance when compared to JULES-INFERNO (e.g., SHAF). In addition, for regions where observations present a positive burnt area trend (TENA, MIDE and SEAS), JULES-INFERNO+HDI presents an opposite trend sign (negative), while JULES-INFERNO is able to capture the
745 positive trend in TENA and SEAS.

It is known that there is an increase in the frequency of large and severe fires in Continental United States of America (Goss et al., 2020; Williams et al., 2019; Abatzoglou and Williams, 2016), as well as boreal regions (Canada and Alaska) (Kasischke and Turetsky, 2006; Stocks et al., 2002; Veraverbeke et al., 2017), leading to observed increases in burnt area, with fire activity having a strong relationship with fire weather in these regions.

750 However, INFERNO has been developed for Earth System Modelling resolutions and timescales. While ESMS can, in principle, represent fire spread using more detailed fire modules (e.g., Lasslop et al., 2014), INFERNO's current implementation relies on the average burnt area for each PFT ($\overline{BA_{PFT}}$) formulation in Eq. 5. This approach does not capture the fine-scale dynamics that drive large and severe fires, meaning that INFERNO may underestimate both the occurrence and extent of extreme fire events. Even in models that explicitly represent fire spread, the ability to resolve these dynamics is constrained
755 by the spatial and temporal resolution at which the ESM is run. Consequently, regions where fire regimes are dominated by large and severe fires may experience a negative bias in predicted burnt areas and fire emissions, as well as in their simulated response to a changing climate.

4.2 Model limitations and known issues

The use of socio-economic factors in INFERNO reduces the inter-annual variability of burnt area for most of the fire regions
760 (Figure 11). While this improves INFERNO performance over regions such as TENA and CEAM, it results in a reduction

in the variability overall, reducing the ability of the model to represent the burnt area regions that are characterised by high inter-annual variability, namely, BONA, BOAS, AUST, CEAS, SHSA and NHSA. Although this further exacerbates the under-estimation of inter-annual variability in JULES-INFERNO+HDI, it must be noted that the control model - JULES-INFERNO - despite having a larger inter-annual variability, also underperforms in this aspect compared to observations.

765 Although socio-economic factors are included in JULES-INFERNO+HDI, the HDI dataset provides information mainly at a national level. To improve the impact of socioeconomic activities on fire at a regional level, it would be beneficial to use data capturing the HDI changes at a sub-national administrative level. Furthermore, it should be highlighted that the HDI does not account for the different implementation of fire management practices and government policies at the regional level.

770 Another limitation in the representation of fires in INFERNO is the lack of a peat-burning in the model simulations described here. The work of Teixeira et al. (2021) highlights that this could be responsible for the negative bias over equatorial Asia and boreal regions where peatland fires represent a significant amount of burnt area and biomass burning emissions. Recent developments from Blackford et al. (2024) could significantly improve the model performance over these regions and help to reduce the burnt area bias both at regional and global scales.

775 It should be noted that, INFERNO represents burnt area as the average per fire for each PFT, decoupling fire spread from localised influences such as wind, weather variability, and topography. This allows the model to capture broad-scale, climate and vegetation-driven fire dynamics through PFT-specific flammability and fire occurrence metrics. However, sub-grid heterogeneity in terrain or meteorology is not explicitly resolved, which may lead to under-representation of local-scale fire spread and associated impacts.

780 In addition, biases in the underlying vegetation can significantly impact modelled burnt area. For example, (Forkel et al., 2019) demonstrated that many global fire models exhibit widespread shortcomings in capturing the sensitivity of burnt area to vegetation characteristics such as leaf area index and plant productivity. Such biases can propagate through modelled fire dynamics, leading to systematic over or underestimation of burnt areas, particularly in regions where vegetation structure strongly regulates fuel availability. The work by Teixeira et al. (2021) is a good example of this. In their work, the authors show that although the burnt area fraction over Africa is well represented, there is a large (50 %) underestimation of the fires
785 in the northern African region (NHAF). This underestimation is attributed to the Saharan bare soil extending too far south, causing a lack of grassland in the Sahel region, which is a result of precipitation deficits associated with errors in the position and intensity of monsoon systems (Sellar et al., 2019b; Williams et al., 2018).

790 Despite the improvements introduced by Burton et al. (2019) to JULES, including fire-vegetation interactions, there are still a number of regions that show significant vegetation biases, which in turn affect the performance of INFERNO. For example, (Burton et al., 2019) shows that JULES vegetation has few needle-leaf trees across the boreal regions compared to observations.

795 These results highlight that the high burnt area variability in these regions may result from a mechanism not currently represented in INFERNO. For example, Kirillina et al. (2020); Andela and Van Der Werf (2014) show how changes in areas with increasing anthropogenic alteration, such as agricultural systems, and changes in fire management practices and government policies, often lead to shifts in peak fire activity for regions such as India and southwest Russia. They also show that the widespread adoption of Aboriginal fire management with increased prescribed burning has curbed the frequency of large fires

over a broad region in Australia. Similarly, the dominant spatial and temporal variability in the burnt area for Southern Europe and North Africa (Chergui et al., 2018), as well as South America (Chuvienco et al., 2021), is known to be driven by shifts in the amounts of fuel and continuity imposed by changes in socioeconomic drivers.

4.3 Concluding remarks

800 Socio-economic policies on management and control of fire play a major role in controlling fire ignition and suppression (Nikolakis and Roberts, 2022; Ford et al., 2021; Jacobson et al., 2022; Carreiras et al., 2014; Mourão and Martinho, 2014). This is especially important in the context of future climate projections (Pivello et al., 2021; Duane et al., 2019; Gillson et al., 2019; Paveglio et al., 2018). It is only with the understanding of the expected impact of climate change that adaptation and mitigation policies can be developed with the aim of protecting infrastructure and ecosystems from fire hazards. This shows the
805 importance of representing socio-economic controls on fire when modelling future projections in an Earth System Modelling context.

This study shows that including a parametrisation for socio-economic impacts on fire based on HDI in INFERNO provides a simple, linear approximation of these effects on fire ignition and suppression. This approach improves the representation of regions where JULES-INFERNO exhibits large positive biases, such as TENA, CEAM, SHSA, EURO, and MIDE. However,
810 the improvements are regionally variable, and in some areas the inclusion of HDI can reduce variability or exacerbate negative biases, highlighting the trade-offs associated with this simplified socio-economic parametrisation.

Introducing socio-economic factors into INFERNO reduces compensating biases and improves the modelled burnt area trends in comparison with observations in several regions. In particular, the results here show that representing the socio-economic factors through HDI in the representation of fires in ESMs can help simulate the burnt area mean state more real-
815 istically, quantify recent past trends, and the relative importance of socio-economic versus climatic and population drivers at regional scales. However, these improvements are not uniform, and in some regions the inclusion of HDI can dampen variability or amplify existing negative biases, highlighting the limitations of a simple, globally applied socio-economic proxy.

The improvements seen in JULES-INFERNO+HDI for regions such as TENA, NHAF, and SHAF contribute more strongly to global metrics than the reduced performance observed in regions such as CEAM, NHSA, SHSA, EURO, and MIDE. How-
820 ever, in regions such as BOAS, CEADS, SEAS, EQAS, and AUST, both model configurations underperform in representing variability, and the differences between the modelled and observed standard deviation (STD / STD_{GFED4s}) remain relatively small compared to the observed variability (e.g., differences between JULES-INFERNO and JULES-INFERNO+HDI STD / STD_{GFED4s} are generally smaller than 15 %). These results highlight that, while the inclusion of HDI can improve global-level metrics, regional performance remains heterogeneous and some limitations in capturing sub-annual variability persist.

825 For regions such as NHSA, BOAS, CEAS, and AUST there is an increase in the bias introduced. However, it should be noted that these model limitations and known issues are discussed in Section 4.2 where the limitations of the model are highlighted and related to fire mechanisms that are not represented in INFERNO, or bias in the underlying vegetation model causing impacts on the modelling of fires. In future work, incorporating processes such as fire intensity, fuel type, seasonal timing, and

vertical extent would be required for a more mechanistic representation of emission injection and atmospheric transport, and
830 may lead to improvements in the model.

Finally, the recent work by Perkins et al. (2024), proposes an alternative methodology for representing human influences in INFERNO. While the WHAM! framework introduced by Perkins et al. (2024) offers a more comprehensive approach, its increased complexity may present challenges in implementation within an ESM context. In contrast, the method presented in this study is straightforward and is directly applicable within the existing modelling framework of JULES-INFERNO for ESM.

835 Considering this, we recommend that socio-economic factors should be included in all fire modelling studies at both global and regional scales, particularly when considering future climate change scenarios. This work will form the basis of a future study on understanding the impact of fires in the Earth System when considering future climate change scenarios.

Appendix A

A1 Burnt area evaluation

840 To analyse the model performance, we calculated the following statistical and error measures, relative to the observed (GFED4s) and modelled (JULES-INFERNO and JULES-INFERNO+HDI) burnt area:

- Deviation of the modelled data in relation to observed values:

$$\phi'_i = \phi_i - \phi_{i,obs} \quad (A1)$$

- Bias, which represents the mean deviation of the modelled data in relation to the observed values.

845

$$Bias = \frac{1}{N} \sum_{i=1}^N \phi'_i \quad (A2)$$

- The Root Mean Square Error.

$$RMSE = \sqrt{\frac{\sum_{i=1}^N (\phi_i - \phi_{i,obs})^2}{N}} \quad (A3)$$

- The Root Mean Square Error after the removal of a constant bias.

$$RMSE_{UB} = \sqrt{\frac{\sum_{i=1}^N [(\phi_i - \bar{\phi}) - (\phi_{i,obs} - \bar{\phi}_{obs})]^2}{N}} \quad (A4)$$

850 – Standard deviation for the modelled - equation A5 - and observed - equation A6 - data.

$$STD = \sqrt{\frac{\sum_{i=1}^N (\phi_i - \bar{\phi})^2}{N}} \quad (A5)$$

$$STD_{obs} = \sqrt{\frac{\sum_{i=1}^N (\phi_{i,obs} - \bar{\phi}_{obs})^2}{N}} \quad (A6)$$

were i is the temporal index and N is the number of elements of ϕ considered, and $\bar{\phi}$ is the constant bias.

Considering these statistics, a perfect simulation would have the following criteria:

855 – RMSE = 0

- $RMSE_{UB} = 0$
- $bias = 0$
- Pearson correlation = 100 %
- $STD / STD_{GFED4s} = 1$
- $RMSE / STD_{GFED4s} = 0$
- $RMSE_{UB} / STD_{GFED4s} = 0$

860

A2 Figures and Tables

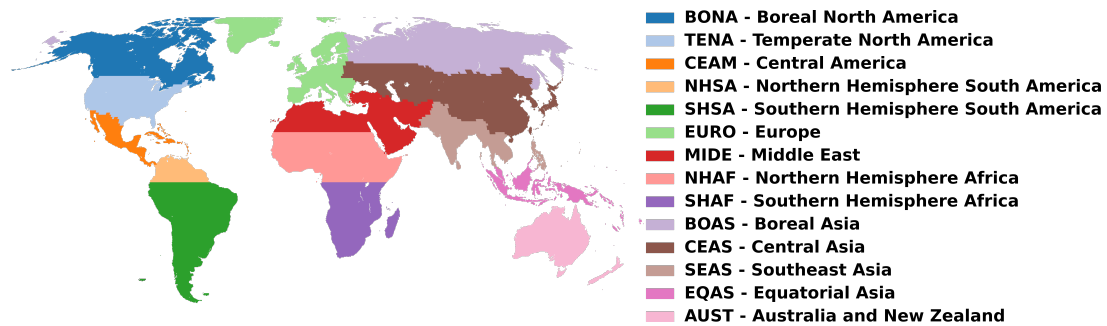


Figure A1. Basis regions, as defined in the GFED4s dataset (Giglio et al., 2013).

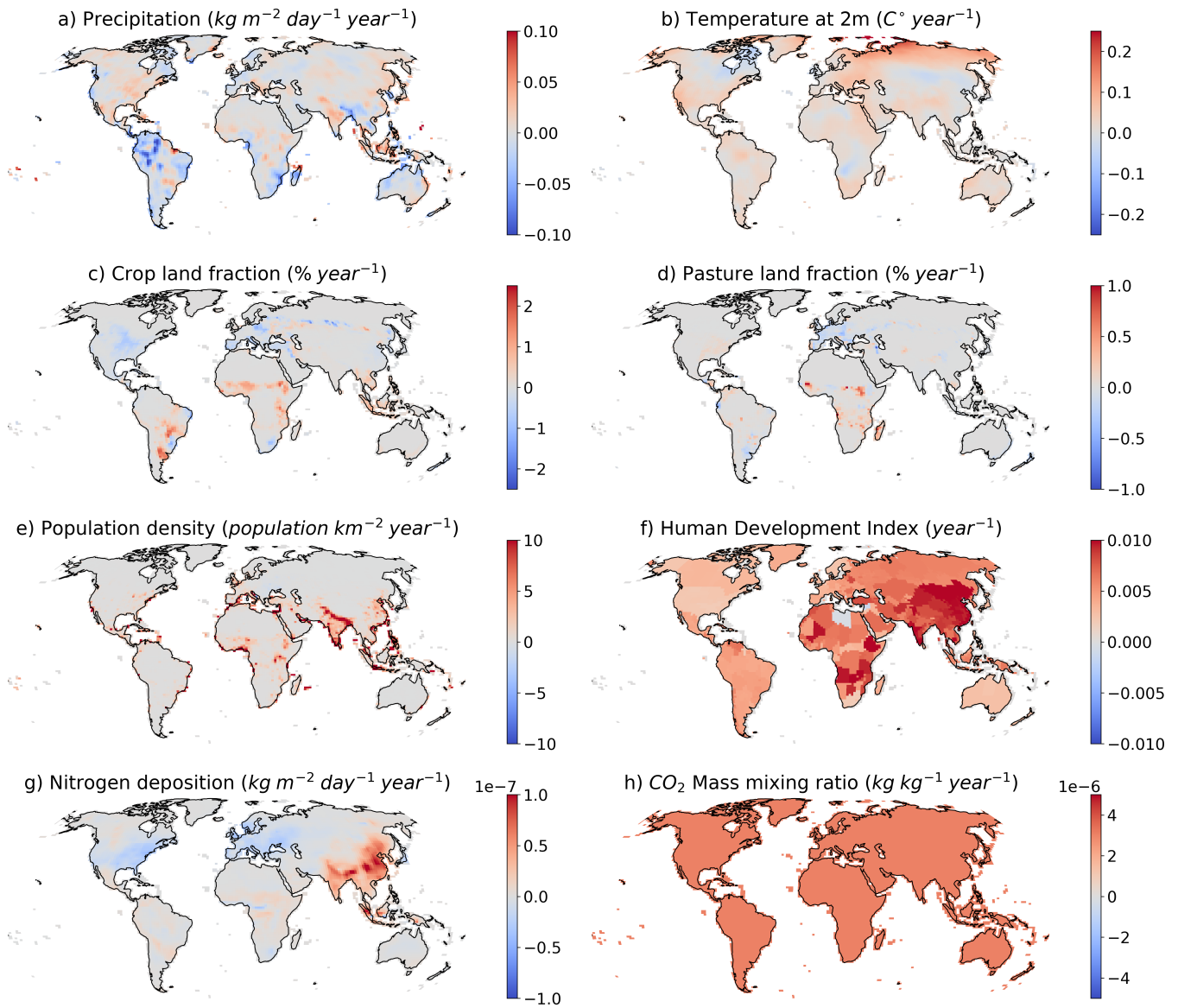


Figure A2. Trends calculated between the period 1997 - 2016 for JULES-ES external forcing variables a) Precipitation ($kg\ m^{-2}\ day^{-1}\ year^{-1}$), b) Temperature at 2 m ($C^{\circ}\ year^{-1}$), c) Crop land fraction ($\% year^{-1}$), d) Pasture land fraction ($\% year^{-1}$), e) Population density ($population\ km^{-2}\ year^{-1}$), f) Human Development Index ($year^{-1}$), g) Nitrogen deposition ($kg\ m^{-2}\ day^{-1}\ year^{-1}$), and h) Carbon Dioxide mixing ratio ($kg\ m^{-2}\ day^{-1}\ year^{-1}$).

Table A1. Annual burnt area statistics for the different GFED4s fire regions (Giglio et al., 2013).

	GLOBAL	BONA	TENA	CEAM	NHSA	SHSA	EURO	MIDE	NHAF	SHAF	BOAS	CEAS	SEAS	EQAS	AUST
GFED4s															
Mean BA (Mha)	476.43	2.47	2.33	1.68	4.94	25.85	0.86	0.90	152.24	155.85	7.65	19.22	12.85	0.70	39.88
Trend (Mha year ⁻¹)	-6.77	0.05	0.05	-0.04	-0.05	-0.44	-0.03	0.02	-2.20	-0.54	-0.03	-0.37	0.18	-0.02	-1.09
Mean BA (Mha)	442.08	4.09	19.53	6.08	1.97	75.09	3.09	4.78	136.30	85.29	10.21	14.88	13.89	0.11	31.43
Trend (Mha year ⁻¹)	-2.24	-0.02	0.09	0.00	0.02	-0.13	0.01	-0.04	-1.26	-0.14	0.05	0.10	0.08	0.00	0.02
Bias (Mha)	-34.35	1.62	17.21	4.40	-2.97	49.24	2.23	3.88	-15.94	-70.56	2.56	-4.34	1.03	-0.59	-8.46
Relative bias (%)	-7.21	65.27	739.79	262.03	-60.16	190.48	258.82	432.09	-10.47	-45.27	33.48	-22.58	8.05	-84.81	-21.20
RMSE	47.28	2.10	17.38	4.52	3.07	49.85	2.27	3.92	20.23	71.17	4.26	7.16	1.91	0.89	21.60
RMSE _{UE}	32.50	1.35	2.43	1.00	0.77	7.78	0.45	0.55	12.45	9.31	3.40	5.69	1.60	0.66	19.87
Correlation	0.753	0.002	0.816	0.548	0.583	0.479	0.153	-0.260	0.777	0.468	0.267	0.138	0.688	0.798	0.870
STD / STD _{GFED4s}	0.52	0.23	4.41	0.84	0.67	0.77	0.68	0.81	0.64	0.43	0.21	0.20	0.50	0.10	0.18
RMSE / STD _{GFED4s}	2.14	1.64	29.46	4.28	3.41	6.18	6.54	11.77	1.35	7.07	1.21	1.33	0.99	1.32	1.14
RMSE _{UE} / STD _{GFED4s}	1.47	1.05	4.13	0.95	0.86	0.96	1.28	1.65	0.83	0.92	0.97	1.06	0.83	0.98	1.05
Mean BA (Mha)	278.90	0.52	3.40	2.13	0.29	25.58	0.45	3.09	125.08	69.65	3.43	6.50	9.56	0.03	7.14
Trend (Mha year ⁻¹)	-7.58	-0.02	-0.07	-0.05	-0.01	-0.64	-0.01	-0.11	-2.71	-1.94	-0.09	-0.22	-0.19	0.00	-0.20
Bias (Mha)	-197.52	-1.95	1.07	0.45	-4.65	-0.27	-0.41	2.19	-27.16	-86.20	-4.22	-12.72	-3.30	-0.67	-32.74
Relative bias (%)	-41.46	-78.97	46.08	26.56	-94.14	-1.05	-48.14	244.36	-17.84	-55.31	-55.17	-66.19	-25.64	-95.88	-82.09
RMSE	198.99	2.37	1.34	1.03	4.73	7.87	0.53	2.35	29.19	87.11	5.44	13.80	4.23	0.97	38.14
RMSE _{UE}	24.12	1.34	0.81	0.92	0.89	7.87	0.33	0.84	10.70	12.60	3.44	5.36	2.66	0.70	19.57
Correlation	0.842	-0.141	0.391	0.614	0.641	0.377	0.732	-0.248	0.863	0.390	0.231	0.370	-0.663	0.850	0.216
STD / STD _{GFED4s}	0.42	0.04	1.04	0.34	0.09	0.28	0.15	0.63	0.68	0.48	0.09	0.10	0.40	0.03	0.06
RMSE / STD _{GFED4s}	9.00	1.85	2.28	0.97	5.26	0.98	1.53	7.06	1.95	8.65	1.55	2.57	2.20	1.44	2.01
RMSE _{UE} / STD _{GFED4s}	1.09	1.04	1.37	0.88	0.98	0.98	0.96	2.53	0.71	1.25	0.98	1.00	1.38	1.04	1.03
JULES-INFERNO+HDI															

Table A2. JULES-INFERNO burnt area trends ($Mha\ year^{-1}$) for the different GFED4s fire regions (Giglio et al., 2013) from the model sensitivity experiments. The standard error of the estimated trend, under the assumption of residual normality is shown in brackets. Values in bold represent trends that are significantly different from zero at the 95 % confidence level.

	JULES-INFERNO									
	GFED4s	control	1990 control	clim	tas	ppn	lu	Ndep	pop	CO2
GLOBAL	-6.77 (9.8e-01)	-2.24 (5.1e-01)	-0.26 (5.3e-01)	-0.06 (5.5e-01)	0.42 (4.7e-01)	-0.92 (4.9e-01)	-1.01 (5.4e-01)	-0.27 (5.3e-01)	-1.45 (4.7e-01)	-0.32 (6.0e-01)
BONA	0.048 (5.2e-02)	-0.017 (1.3e-02)	0.002 (1.9e-03)	-0.023 (1.2e-02)	0.008 (1.5e-02)	-0.007 (7.7e-03)	0.003 (1.8e-03)	0.002 (1.8e-03)	0.005 (1.8e-03)	0.003 (1.8e-03)
TENA	0.05 (2.4e-02)	0.08 (1.2e-01)	0.01 (1.1e-02)	0.04 (1.1e-01)	0.08 (2.9e-02)	-0.05 (6.7e-02)	0.05 (9.3e-03)	0.01 (1.2e-02)	-0.02 (1.1e-02)	0.05 (1.2e-02)
CEAM	-0.07 (4.3e-02)	0.000 (3.9e-02)	0.01 (9.1e-03)	0.03 (4.2e-02)	0.03 (1.4e-02)	0.02 (3.9e-02)	0.01 (8.6e-03)	0.01 (9.1e-03)	-0.01 (8.9e-03)	0.01 (9.2e-03)
NHSA	-0.05 (3.7e-02)	0.02 (2.6e-02)	0.001 (7.5e-04)	0.03 (2.6e-02)	0.01 (6.0e-03)	0.02 (1.8e-02)	0.002 (9.4e-04)	0.001 (7.5e-04)	-0.002 (7.9e-04)	-0.005 (1.5e-03)
SHSA	-0.41 (3.3e-01)	-0.12 (2.6e-01)	0.04 (8.3e-02)	0.14 (2.8e-01)	0.27 (7.3e-02)	-0.17 (2.2e-01)	-0.17 (8.6e-02)	0.04 (8.3e-02)	-0.005 (8.2e-02)	-0.01 (9.4e-02)
EURO	-0.03 (1.4e-02)	0.01 (1.1e-02)	0.001 (3.5e-03)	-0.001 (1.0e-02)	0.01 (3.9e-03)	-0.01 (7.4e-03)	0.01 (2.0e-03)	0.0005 (3.4e-03)	-0.01 (3.5e-03)	0.003 (3.2e-03)
MIDE	0.02 (1.4e-02)	-0.04 (1.2e-02)	0.001 (7.8e-03)	-0.02 (1.3e-02)	0.01 (9.9e-03)	-0.04 (1.2e-02)	0.001 (7.8e-03)	-0.003 (9.4e-03)	-0.04 (6.9e-03)	0.01 (7.6e-03)
NHAF	-2.24 (6.1e-01)	-1.35 (5.1e-01)	0.04 (3.1e-01)	0.15 (5.2e-01)	0.14 (3.2e-01)	-0.10 (5.3e-01)	-0.35 (3.3e-01)	0.04 (3.1e-01)	-0.55 (2.7e-01)	0.06 (3.2e-01)
SHAF	-0.53 (4.1e-01)	-0.15 (1.9e-01)	0.17 (1.4e-01)	0.27 (1.9e-01)	0.15 (1.4e-01)	0.15 (1.4e-01)	0.01 (1.4e-01)	0.17 (1.4e-01)	-0.18 (1.4e-01)	0.04 (1.6e-01)
BOAS	-0.07 (1.4e-01)	0.05 (3.1e-02)	0.01 (4.2e-03)	0.04 (2.9e-02)	0.06 (2.1e-02)	0.03 (1.5e-02)	0.02 (3.9e-03)	0.01 (4.0e-03)	-0.01 (3.9e-03)	0.01 (4.1e-03)
CEAS	-0.35 (2.2e-01)	0.10 (4.6e-02)	0.01 (7.3e-03)	0.08 (5.0e-02)	0.05 (2.4e-02)	0.03 (2.6e-02)	0.02 (8.0e-03)	0.01 (6.7e-03)	0.01 (7.5e-03)	0.03 (9.2e-03)
SEAS	0.17 (7.9e-02)	0.07 (4.4e-02)	0.02 (8.6e-03)	0.06 (4.3e-02)	0.03 (2.4e-02)	0.05 (3.3e-02)	0.0004 (7.6e-03)	0.02 (7.9e-03)	0.04 (8.4e-03)	0.02 (9.0e-03)
EQAS	-0.04 (2.8e-02)	-0.0027 (3.0e-03)	0.0004 (2.2e-04)	-0.0020 (3.3e-03)	0.0003 (3.3e-04)	-0.0017 (2.5e-03)	-0.0001 (2.2e-04)	0.0004 (2.2e-04)	-0.0002 (2.0e-04)	0.0003 (2.2e-04)
AUST	-1.02 (7.8e-01)	0.01 (1.4e-01)	0.07 (5.6e-02)	-0.05 (1.5e-01)	0.12 (5.9e-02)	-0.12 (1.3e-01)	0.07 (5.6e-02)	0.06 (5.6e-02)	0.07 (5.6e-02)	0.11 (5.8e-02)

Table A3. JULES-INFERNO+HDI burnt area trends ($Mha\ year^{-1}$) for the different GFED4s fire regions (Giglio et al., 2013) from the model sensitivity experiments. The standard error of the estimated trend, under the assumption of residual normality is shown in brackets. Values in bold represent trends that are significantly different from zero at the 95 % confidence level.

	JULES-INFERNO+HDI										
	GFED4s	control	1990 control	clim	tas	ppn	lu	Ndep	pop	CO2	HDI
GLOBAL	-6.77 (9.8e-01)	0.15 (2.2e-02)	-7.58 (4.2e-01)	0.24 (5.5e-01)	0.58 (2.1e-01)	-0.60 (5.5e-01)	-0.97 (2.0e-02)	0.16 (2.3e-02)	-0.91 (1.9e-02)	0.02 (2.3e-02)	-4.27 (9.8e-02)
BONA	0.048 (5.2e-02)	-0.0002 (8.2e-05)	-0.023 (2.1e-03)	-0.006 (2.8e-03)	0.002 (2.8e-03)	-0.003 (1.9e-03)	-0.0002 (8.2e-05)	-0.0002 (8.3e-05)	-0.0003 (8.0e-05)	-0.0001 (8.1e-05)	-0.021 (6.6e-04)
TENA	0.05 (2.4e-02)	0.00 (7.1e-04)	-0.07 (2.8e-02)	0.001 (3.6e-02)	0.02 (8.3e-03)	-0.02 (2.2e-02)	0.01 (8.4e-04)	0.000 (8.3e-04)	-0.01 (7.9e-04)	0.004 (8.0e-04)	-0.05 (1.7e-03)
CEAM	-0.07 (4.3e-02)	-0.01 (1.5e-03)	-0.04 (1.6e-02)	0.02 (2.3e-02)	0.01 (3.8e-03)	0.01 (2.3e-02)	-0.01 (1.6e-03)	-0.01 (1.5e-03)	-0.02 (1.4e-03)	-0.01 (1.7e-03)	-0.05 (1.5e-03)
NHSA	-0.05 (3.7e-02)	0.000 (1.2e-04)	-0.01 (3.7e-03)	0.01 (5.8e-03)	0.003 (1.5e-03)	0.003 (3.9e-03)	-0.001 (1.0e-04)	-0.0005 (1.2e-04)	-0.001 (1.3e-04)	-0.004 (1.5e-04)	-0.01 (4.7e-04)
SHSA	-0.41 (3.3e-01)	0.05 (2.1e-02)	-0.62 (9.7e-02)	0.11 (1.4e-01)	0.14 (3.1e-02)	-0.09 (1.1e-01)	-0.06 (2.4e-02)	0.05 (2.1e-02)	0.02 (2.1e-02)	-0.002 (2.3e-02)	-0.34 (1.9e-02)
EURO	-0.03 (1.4e-02)	-0.0002 (1.2e-04)	-0.01 (2.4e-03)	-0.0001 (3.1e-03)	0.004 (9.2e-04)	-0.003 (2.1e-03)	0.004 (1.3e-04)	-0.0002 (1.2e-04)	-0.002 (1.5e-04)	0.001 (1.3e-04)	-0.01 (4.4e-04)
MIDE	0.02 (1.4e-02)	0.02 (3.0e-03)	-0.11 (8.9e-03)	-0.004 (1.5e-02)	0.03 (5.7e-03)	-0.02 (1.5e-02)	0.02 (2.6e-03)	0.02 (4.4e-03)	-0.03 (3.0e-03)	0.04 (2.5e-03)	-0.09 (4.8e-03)
NHAF	-2.24 (6.1e-01)	0.14 (7.1e-02)	-2.78 (5.5e-01)	0.50 (6.3e-01)	0.33 (1.2e-01)	0.12 (6.0e-01)	-0.57 (6.6e-02)	0.14 (7.2e-02)	-0.31 (7.3e-02)	0.24 (7.6e-02)	-0.73 (9.0e-02)
SHAF	-0.53 (4.1e-01)	0.29 (1.3e-01)	-1.92 (2.1e-01)	0.28 (2.2e-01)	0.24 (1.5e-01)	0.05 (1.5e-01)	0.03 (1.3e-01)	0.30 (1.3e-01)	-0.09 (1.3e-01)	0.07 (1.3e-01)	-1.48 (1.7e-01)
BOAS	-0.07 (1.4e-01)	0.001 (9.0e-04)	-0.10 (1.3e-02)	0.02 (1.2e-02)	0.02 (6.2e-03)	0.02 (6.8e-03)	0.005 (8.8e-04)	0.002 (9.3e-04)	-0.002 (8.7e-04)	0.003 (9.1e-04)	-0.12 (5.0e-03)
CEAS	-0.35 (2.2e-01)	0.01 (5.0e-03)	-0.23 (2.3e-02)	0.04 (3.1e-02)	0.04 (1.2e-02)	0.02 (1.7e-02)	0.02 (4.5e-03)	0.02 (4.4e-03)	0.01 (5.0e-03)	0.03 (4.9e-03)	-0.23 (8.1e-03)
SEAS	0.17 (7.9e-02)	0.02 (7.7e-03)	-0.19 (3.5e-02)	0.08 (4.5e-02)	0.04 (2.5e-02)	0.06 (3.7e-02)	0.004 (7.4e-03)	0.02 (7.0e-03)	0.04 (7.6e-03)	0.02 (7.6e-03)	-0.19 (7.7e-03)
EQAS	-0.04 (2.8e-02)	0.0003 (1.1e-04)	-0.0014 (8.3e-04)	-0.0005 (1.3e-03)	0.0003 (1.3e-04)	-0.0005 (9.9e-04)	-0.0001 (1.1e-04)	0.0003 (1.1e-04)	0.0001 (9.5e-05)	0.0002 (1.1e-04)	-0.001 (1.0e-04)
AUST	-1.02 (7.8e-01)	0.03 (1.9e-02)	-0.19 (5.3e-02)	0.01 (7.2e-02)	0.06 (2.5e-02)	-0.01 (6.5e-02)	0.03 (1.9e-02)	0.03 (1.9e-02)	0.03 (1.9e-02)	0.04 (1.9e-02)	-0.17 (1.6e-02)

Code and data availability. Both the model code and the files for running it are available from the Met Office Science Repository Service: <https://code.metoffice.gov.uk/> (last access: 26 July 2023). Registration is required, and code is freely available subject to completion of a software license.

Details of the simulations performed: JULES simulations are compiled and run in suites developed using the Rose suite engine (MetOffice, 2022) and scheduled using the cylc workflow engine (Oliver et al., 2019). Both Rose and cylc are available under v3 of the GNU General Public License (GPL). In this framework, the suite contains the information required to extract and build the code as well as configure and run the simulations. Each suite is labelled with a unique identifier and is held in the same revision-controlled repository service in which we hold and develop the model code. This means that these suites are available to any licensed of JULES under the following suite IDs:

- JULES-INFERNO: u-by849
- JULES-INFERNO+HDI: u-by851

For JULES-INFERNO sensitivity experiments:

- 1990 control: u-co594
- clim: u-cs067
- tas: u-cs068
- ppn: u-cs069
- lu: u-cr440
- Ndep: u-cr441

- pop: u-cr442
- CO₂: u-cr443

For JULES-INFERNO+HDI sensitivity experiments:

- 1990 control: u-ct759
- clim: u-cs070
- tas: u-cs071
- ppn: u-cs072
- lu: u-cr447
- Ndep: u-cr448
- pop: u-cr449
- CO₂: u-cr450
- HDI: u-cn957

Author contributions. JCMT led the writing of the paper and model development. All co-authors contributed to the simulation design, writing sections, performing evaluation and reviewing drafts of the paper.

Competing interests. At least one of the (co-)authors is a member of the editorial board of Earth System Dynamics.

895 *Acknowledgements.* We would especially like to thank those who have contributed to the development of INFERNO, with a special thanks to Stéphane Mangeon for taking the first steps to develop INFERNO and the observational community, who have developed numerous datasets used in this paper to help evaluate the model.

900 This research and JCMT, CB, GAF, FMOC, RAB were supported by the Met Office Hadley Centre Climate Programme funded by DSIT. JCMT, GAF, and FMOC were also supported by the Horizon 2020 Framework Programme (CRESCENDO, grant no. 779366) and the Earth System Models for the Future (ESM2025, grant no. 101003536). CB was funded by the Met Office Climate Science for Service Partnership (CSSP) Brazil project which is supported by the Department for Science, Innovation & Technology (DSIT). DIK was supported by the Natural Environment Research Council as part of the LTSM2 TerraFIRMA project. AV was funded via the Leverhulme Centre for Wildfires, Environment and Society through the Leverhulme Trust, grant no. RC-2018-023.

References

- 905 Abatzoglou, J. T. and Williams, A. P.: Impact of anthropogenic climate change on wildfire across western US forests, *Proceedings of the National Academy of Sciences*, 113, 11 770–11 775, 2016.
- Abatzoglou, J. T., Williams, A. P., Boschetti, L., Zubkova, M., and Kolden, C. A.: Global patterns of interannual climate–fire relationships, *Global change biology*, 24, 5164–5175, 2018.
- Andela, N. and Van Der Werf, G. R.: Recent trends in African fires driven by cropland expansion and El Nino to La Nina transition, *Nature*
910 *Climate Change*, 4, 791–795, 2014.
- Andela, N., Morton, D. C., Giglio, L., Chen, Y., van der Werf, G. R., Kasibhatla, P. S., DeFries, R. S., Collatz, G., Hantson, S., Kloster, S., et al.: A human-driven decline in global burned area, *Science*, 356, 1356–1362, 2017.
- Andela, N., Morton, D. C., Giglio, L., Paugam, R., Chen, Y., Hantson, S., Van Der Werf, G. R., and Randerson, J. T.: The Global Fire Atlas of individual fire size, duration, speed and direction, *Earth System Science Data*, 11, 529–552, 2019.
- 915 Aragão, L. E., Anderson, L. O., Fonseca, M. G., Rosan, T. M., Vedovato, L. B., Wagner, F. H., Silva, C. V., Silva Junior, C. H., Arai, E., Aguiar, A. P., et al.: 21st Century drought-related fires counteract the decline of Amazon deforestation carbon emissions, *Nature communications*, 9, 536, 2018.
- Benjamini, Y. and Yekutieli, D.: The control of the false discovery rate in multiple testing under dependency, *Annals of statistics*, pp. 1165–1188, 2001.
- 920 Best, M. J., Pryor, M., Clark, D. B., Rooney, G. G., Essery, R. L. H., Ménard, C. B., Edwards, J. M., Hendry, M. A., Porson, A., Gedney, N., Mercado, L. M., Sitch, S., Blyth, E., Boucher, O., Cox, P. M., Grimmond, C. S. B., and Harding, R. J.: The Joint UK Land Environment Simulator (JULES), model description – Part 1: Energy and water fluxes, *Geoscientific Model Development*, 4, 677–699, <https://doi.org/10.5194/gmd-4-677-2011>, 2011.
- Bhanojirao, V.: Human development report 1990: review and assessment, *World Development*, 19, 1451–1460, 1991.
- 925 Blackford, K. R., Kasoar, M., Burton, C., Burke, E., Prentice, I. C., and Voulgarakis, A.: INFERNO-peat v1. 0.0: a representation of northern high-latitude peat fires in the JULES-INFERNO global fire model, *Geoscientific Model Development*, 17, 3063–3079, 2024.
- Bowman, D. M., Kolden, C. A., Abatzoglou, J. T., Johnston, F. H., van der Werf, G. R., and Flannigan, M.: Vegetation fires in the Anthropocene, *Nature Reviews Earth & Environment*, 1, 500–515, 2020.
- Burton, C., Betts, R., Cardoso, M., Feldpausch, T. R., Harper, A., Jones, C. D., Kelley, D. I., Robertson, E., and Wiltshire, A.: Representation
930 of fire, land-use change and vegetation dynamics in the Joint UK Land Environment Simulator vn4.9 (JULES), *Geoscientific Model Development*, 12, 179–193, <https://doi.org/10.5194/gmd-12-179-2019>, 2019.
- Burton, C., Betts, R. A., Jones, C. D., Feldpausch, T. R., Cardoso, M., and Anderson, L. O.: El Niño Driven Changes in Global Fire 2015/16, *Frontiers in Earth Science*, 8, 199, <https://doi.org/10.3389/feart.2020.00199>, 2020.
- Burton, C. A., Kelley, D. I., Burke, E., Mathison, C., Jones, C. D., Betts, R. A., Robertson, E., Teixeira, J. C., Cardoso, M., and Anderson,
935 L. O.: Fire weakens land carbon sinks before 1.5° C, *Nature Geoscience*, 17, 1108–1114, 2024.
- Carreiras, M., Ferreira, A. J. D., Valente, S., Fleskens, L., Gonzales-Pelayo, Ó., Rubio, J. L., Stoof, C. R., Coelho, C. O. A., Ferreira, C. S. S., and Ritsema, C. J.: Comparative analysis of policies to deal with wildfire risk, *Land Degradation & Development*, 25, 92–103, 2014.
- Cecil, D.: LIS/OTD 0.5 Degree High Resolution Monthly Climatology (HRMC), <https://doi.org/10.3389/feart.2020.00199>, 2006.
- Chergui, B., Fahd, S., Santos, X., and Pausas, J. G.: Socioeconomic factors drive fire-regime variability in the Mediterranean Basin, *Ecosys-*
940 *tems*, 21, 619–628, 2018.

- Christian, H. J., Blakeslee, R. J., Boccippio, D. J., Boeck, W. L., Buechler, D. E., Driscoll, K. T., Goodman, S. J., Hall, J. M., Koshak, W. J., Mach, D. M., et al.: Global frequency and distribution of lightning as observed from space by the Optical Transient Detector, *Journal of Geophysical Research: Atmospheres*, 108, ACL-4, 2003.
- 945 Chuvieco, E., Pettinari, M. L., Koutsias, N., Forkel, M., Hantson, S., and Turco, M.: Human and climate drivers of global biomass burning variability, *Science of the Total Environment*, 779, 146-361, 2021.
- Clark, D. B., Mercado, L. M., Sitch, S., Jones, C. D., Gedney, N., Best, M. J., Pryor, M., Rooney, G. G., Essery, R. L. H., Blyth, E., Boucher, O., Harding, R. J., Huntingford, C., and Cox, P. M.: The Joint UK Land Environment Simulator (JULES), model description – Part 2: Carbon fluxes and vegetation dynamics, *Geoscientific Model Development*, 4, 701–722, <https://doi.org/10.5194/gmd-4-701-2011>, 2011.
- Cox, P.: Description of the "TRIFFID" Dynamic Global Vegetation Mode, 2001.
- 950 Cox, P., Betts, R., Jones, C., Spall, S., and Totterdell, I.: Acceleration of global warming due to carbon-cycle feedbacks in a coupled climate model, *Nature*, 408, 184–187, 2000.
- Curt, T. and Frejaville, T.: Wildfire policy in Mediterranean France: how far is it efficient and sustainable?, *Risk analysis*, 38, 472–488, 2018.
- Duane, A., Aquilué, N., Canelles, Q., Morán-Ordoñez, A., De Cáceres, M., and Brotons, L.: Adapting prescribed burns to future climate change in Mediterranean landscapes, *Science of the Total Environment*, 677, 68–83, 2019.
- 955 Eyring, V., Bony, S., Meehl, G. A., Senior, C. A., Stevens, B., Stouffer, R. J., and Taylor, K. E.: Overview of the Coupled Model Intercomparison Project Phase 6 (CMIP6) experimental design and organization, *Geoscientific Model Development*, 9, 1937–1958, 2016.
- Field, R., Spessa, A., Aziz, N., Camia, A., Cantin, A., Carr, R., De Groot, W., Dowdy, A., Flannigan, M., Manomaiphiboon, K., et al.: Development of a global fire weather database, *Natural Hazards and Earth System Sciences*, 15, 1407–1423, 2015.
- Ford, A. E., Harrison, S. P., Kountouris, Y., Millington, J. D., Mistry, J., Perkins, O., Rabin, S. S., Rein, G., Schreckenberg, K., Smith, C., 960 et al.: Modelling human-fire interactions: combining alternative perspectives and approaches, *Frontiers in Environmental Science*, p. 418, 2021.
- Forkel, M., Andela, N., Harrison, S. P., Lasslop, G., Van Marle, M., Chuvieco, E., Dorigo, W., Forrest, M., Hantson, S., Heil, A., et al.: Emergent relationships with respect to burned area in global satellite observations and fire-enabled vegetation models, *Biogeosciences*, 16, 57–76, 2019.
- 965 Giglio, L., Randerson, J. T., and Van Der Werf, G. R.: Analysis of daily, monthly, and annual burned area using the fourth-generation global fire emissions database (GFED4), *Journal of Geophysical Research: Biogeosciences*, 118, 317–328, 2013.
- Gillson, L., Whitlock, C., and Humphrey, G.: Resilience and fire management in the Anthropocene, *Ecology and Society*, 24, 2019.
- Goldewijk, K. K., Dekker, S. C., and van Zanden, J. L.: Per-capita estimations of long-term historical land use and the consequences for global change research, *Journal of Land Use Science*, 12, 313–337, <https://doi.org/10.1080/1747423X.2017.1354938>, 2017.
- 970 Goss, M., Swain, D. L., Abatzoglou, J. T., Sarhadi, A., Kolden, C. A., Williams, A. P., and Diffenbaugh, N. S.: Climate change is increasing the likelihood of extreme autumn wildfire conditions across California, *Environmental Research Letters*, 15, 094 016, 2020.
- Haas, O., Prentice, I. C., and Harrison, S. P.: Global environmental controls on wildfire burnt area, size, and intensity, *Environmental Research Letters*, 17, 065 004, 2022.
- Haas, O., Prentice, I., and Harrison, S.: Global wildfires on a changing planet - in review, 2024.
- 975 Harris, I., Jones, P. D., Osborn, T. J., and Lister, D. H.: Updated high-resolution grids of monthly climatic observations—the CRU TS3. 10 Dataset, *International journal of climatology*, 34, 623–642, 2014.
- Hickel, J.: The sustainable development index: Measuring the ecological efficiency of human development in the anthropocene, *Ecological economics*, 167, 106 331, 2020.

- Jacobson, M., Smith, H., Huber-Stearns, H. R., Davis, E. J., Cheng, A. S., and Deak, A.: Comparing social constructions of wildfire risk across media, government, and participatory discourse in a Colorado fireshed, *Journal of Risk Research*, 25, 697–714, 2022.
- Jones, M. W., Abatzoglou, J. T., Veraverbeke, S., Andela, N., Lasslop, G., Forkel, M., Smith, A. J., Burton, C., Betts, R. A., van der Werf, G. R., et al.: Global and regional trends and drivers of fire under climate change, *Reviews of Geophysics*, p. e2020RG000726, 2022.
- Jones, M. W., Kelley, D. I., Burton, C. A., Di Giuseppe, F., Barbosa, M. L. F., Brambleby, E., Hartley, A. J., Lombardi, A., Mataveli, G., McNorton, J. R., et al.: State of wildfires 2023–2024, *Earth System Science Data*, 16, 3601–3685, 2024.
- 985 Kasischke, E. S. and Turetsky, M. R.: Recent changes in the fire regime across the North American boreal region—Spatial and temporal patterns of burning across Canada and Alaska, *Geophysical research letters*, 33, 2006.
- Kelley, D. I., Bistinas, I., Whitley, R., Burton, C., Marthews, T. R., and Dong, N.: How contemporary bioclimatic and human controls change global fire regimes, *Nature Climate Change*, 9, 690–696, 2019.
- Kirillina, K., Shvetsov, E. G., Protopopova, V. V., Thiesmeyer, L., and Yan, W.: Consideration of anthropogenic factors in boreal forest fire regime changes during rapid socio-economic development: case study of forestry districts with increasing burnt area in the Sakha Republic, Russia, *Environmental Research Letters*, 15, 035009, 2020.
- Klauenberg, K., Wübbeler, G., Mickan, B., Harris, P., and Elster, C.: A tutorial on Bayesian normal linear regression, *Metrologia*, 52, 878, 2015.
- Klein Goldewijk, K., Beusen, A., Doelman, J., and Stehfest, E.: Anthropogenic land use estimates for the Holocene – HYDE 3.2, *Earth System Science Data*, 9, 927–953, <https://doi.org/10.5194/essd-9-927-2017>, 2017.
- 995 Kumm, M., Taka, M., and Guillaume, J. H.: Gridded global datasets for gross domestic product and Human Development Index over 1990–2015, *Scientific data*, 5, 180004, 2018.
- Lasslop, G., Thonicke, K., and Kloster, S.: SPITFIRE within the MPI Earth system model: Model development and evaluation, *Journal of Advances in Modeling Earth Systems*, 6, 740–755, 2014.
- 1000 Li, F., Levis, S., and Ward, D.: Quantifying the role of fire in the Earth system—Part 1: Improved global fire modeling in the Community Earth System Model (CESM1), *Biogeosciences*, 10, 2293–2314, 2013.
- Li, F., Song, X., Harrison, S. P., Marlon, J. R., Lin, Z., Leung, L. R., Schwinger, J., Maréchal, V., Wang, S., Ward, D. S., et al.: Evaluation of global fire simulations in CMIP6 Earth system models, *Geoscientific Model Development Discussions*, 2024, 1–37, 2024.
- Mangeon, S., Voulgarakis, A., Gilham, R., Harper, A., Sitch, S., and Folberth, G.: INFERNO: a fire and emissions scheme for the UK Met Office’s Unified Model, *Geoscientific Model Development*, 9, 2685–2700, 2016.
- 1005 Marlon, J. R., Bartlein, P. J., Carcaillet, C., Gavin, D. G., Harrison, S. P., Higuera, P. E., Joos, F., Power, M., and Prentice, I.: Climate and human influences on global biomass burning over the past two millennia, *Nature Geoscience*, 1, 697–702, 2008.
- Mathison, C., Burke, E., Hartley, A. J., Kelley, D. I., Burton, C., Robertson, E., Gedney, N., Williams, K., Wiltshire, A., Ellis, R. J., et al.: Description and Evaluation of the JULES-ES setup for ISIMIP2b, *EGU sphere*, 2022, 1–24, 2022.
- 1010 MetOffice: Rose suite engine, <https://metomi.github.io/rose/doc/html/index.html>, 2022.
- Miranda-Lescano, R., Muinel-Gallo, L., and Roca-Sagalés, O.: Human development and decentralization: The importance of public health expenditure, *Annals of Public and Cooperative Economics*, 94, 191–219, 2023.
- Mourão, P. R. and Martinho, V. D.: The choices of the fire—Debating socioeconomic determinants of the fires observed at Portuguese municipalities, *Forest Policy and Economics*, 43, 29–40, 2014.

- 1015 Nepstad, D., McGrath, D., Stickler, C., Alencar, A., Azevedo, A., Swette, B., Bezerra, T., DiGiano, M., Shimada, J., Seroa da Motta, R., et al.: Slowing Amazon deforestation through public policy and interventions in beef and soy supply chains, *science*, 344, 1118–1123, 2014.
- Nikolakis, W. and Roberts, E.: Wildfire governance in a changing world: Insights for policy learning and policy transfer, *Risk, Hazards & Crisis in Public Policy*, 13, 144–164, 2022.
- 1020 Oliver, H., Shin, M., Matthews, D., Sanders, O., Bartholomew, S., Clark, A., Fitzpatrick, B., van Haren, R., Hut, R., and Drost, N.: Workflow automation for cycling systems, *Computing in Science & Engineering*, 21, 7–21, 2019.
- Pandey, P., Huidobro, G., Lopes, L. F., Ganteaume, A., Ascoli, D., Colaco, C., Xanthopoulos, G., Giannaros, T. M., Gazzard, R., Boustras, G., et al.: A global outlook on increasing wildfire risk: Current policy situation and future pathways, *Trees, Forests and People*, 14, 100431, 2023.
- 1025 Paveglio, T. B., Carroll, M. S., Stasiewicz, A. M., Williams, D. R., and Becker, D. R.: Incorporating social diversity into wildfire management: Proposing “pathways” for fire adaptation, *Forest Science*, 64, 515–532, 2018.
- Pechony, O. and Shindell, D.: Fire parameterization on a global scale, *Journal of Geophysical Research: Atmospheres*, 114, 2009.
- Perkins, O., Kasoar, M., Voulgarakis, A., Smith, C., Mistry, J., and Millington, J. D.: A global behavioural model of human fire use and management: WHAM! v1. 0, *Geoscientific Model Development*, 17, 3993–4016, 2024.
- 1030 Pivello, V. R., Vieira, I., Christianini, A. V., Ribeiro, D. B., da Silva Menezes, L., Berlinck, C. N., Melo, F. P., Marengo, J. A., Tornquist, C. G., Tomas, W. M., et al.: Understanding Brazil’s catastrophic fires: Causes, consequences and policy needed to prevent future tragedies, *Perspectives in Ecology and Conservation*, 19, 233–255, 2021.
- Prestemon, J. P., Butry, D. T., Abt, K. L., and Sutphen, R.: Net benefits of wildfire prevention education efforts, *Forest Science*, 56, 181–192, 2010.
- 1035 Rabin, S. S., Melton, J. R., Lasslop, G., Bachelet, D., Forrest, M., Hantson, S., Kaplan, J. O., Li, F., Mangeon, S., Ward, D. S., et al.: The Fire Modeling Intercomparison Project (FireMIP), phase 1: experimental and analytical protocols with detailed model descriptions, *Geoscientific Model Development*, 10, 1175–1197, 2017.
- Ramdas, A., García Trillos, N., and Cuturi, M.: On wasserstein two-sample testing and related families of nonparametric tests, *Entropy*, 19, 47, 2017.
- 1040 Randerson, J., Chen, Y., Van Der Werf, G., Rogers, B., and Morton, D.: Global burned area and biomass burning emissions from small fires, *Journal of Geophysical Research: Biogeosciences*, 117, 2012.
- Rideout, D. B., Wei, Y., Kirsch, A., and Kernohan, N.: STARFire: Strategic budgeting and planning for wildland fire management, *Park Science*, 32, 34–41, 2017.
- Riley, K. L., Williams, A. P., Urbanski, S. P., Calkin, D. E., Short, K. C., and O’Connor, C. D.: Will landscape fire increase in the future? A systems approach to climate, fire, fuel, and human drivers, *Current Pollution Reports*, 5, 9–24, 2019.
- 1045 Rizzo, L. V. and Rizzo, M. C. F.: Wildfire smoke and health impacts: a narrative review, *Jornal de Pediatria*, 2024.
- Roy, A., Dutta, T., Li, Y., and Dong, X.: Human development at the cost of the environment?—an application of planetary pressures—adjusted human development index in the lens of planetary boundaries, *Environmental Science and Pollution Research*, 30, 32 383–32 405, 2023.
- Sellar, A. A., Jones, C. G., Mulcahy, J. P., Tang, Y., Yool, A., Wiltshire, A., O’Connor, F. M., Stringer, M., Hill, R., Palmieri, J., Woodward, S., de Mora, L., Kuhlbrodt, T., Rumbold, S. T., Kelley, D. I., Ellis, R., Johnson, C. E., Walton, J., Abraham, N. L., Andrews, M. B., Andrews, T., Archibald, A. T., Berthou, S., Burke, E., Blockley, E., Carslaw, K., Dalvi, M., Edwards, J., Folberth, G. A., Gedney, N., Griffiths, P. T., Harper, A. B., Hendry, M. A., Hewitt, A. J., Johnson, B., Jones, A., Jones, C. D., Keeble, J., Liddicoat, S., Morgenstern,

- O., Parker, R. J., Predoi, V., Robertson, E., Siahhan, A., Smith, R. S., Swaminathan, R., Woodhouse, M. T., Zeng, G., and Zerroukat, M.: UKESM1: Description and Evaluation of the U.K. Earth System Model, *Journal of Advances in Modeling Earth Systems*, 11, 4513–4558, <https://doi.org/https://doi.org/10.1029/2019MS001739>, 2019a.
- 1055 Sellar, A. A., Jones, C. G., Mulcahy, J. P., Tang, Y., Yool, A., Wiltshire, A., O’connor, F. M., Stringer, M., Hill, R., Palmieri, J., et al.: UKESM1: Description and evaluation of the UK Earth System Model, *Journal of Advances in Modeling Earth Systems*, 11, 4513–4558, 2019b.
- Silva Junior, C. H., Pessôa, A. C., Carvalho, N. S., Reis, J. B., Anderson, L. O., and Aragão, L. E.: The Brazilian Amazon deforestation rate in 2020 is the greatest of the decade, *Nature Ecology & Evolution*, 5, 144–145, 2021.
- 1060 Stocks, B., Mason, J., Todd, J., Bosch, E., Wotton, B., Amiro, B., Flannigan, M., Hirsch, K., Logan, K., Martell, D., et al.: Large forest fires in Canada, 1959–1997, *Journal of Geophysical Research: Atmospheres*, 107, FFR–5, 2002.
- Sullivan, A., Kurvits, T., and E., B.: Spreading like Wildfire – The Rising Threat of Extraordinary Landscape Fires., United Nations Environment Programme (UNEP), 2022.
- 1065 Teckentrup, L., Harrison, S. P., Hantson, S., Heil, A., Melton, J. R., Forrest, M., Li, F., Yue, C., Arneth, A., Hickler, T., et al.: Response of simulated burned area to historical changes in environmental and anthropogenic factors: a comparison of seven fire models, *Biogeosciences*, 16, 3883–3910, 2019.
- Teixeira, J. C., Folberth, G. A., O’Connor, F. M., Unger, N., and Voulgarakis, A.: Coupling interactive fire with atmospheric composition and climate in the UK Earth System Model, *Geoscientific Model Development*, 14, 6515–6539, 2021.
- 1070 Türe, C.: A methodology to analyse the relations of ecological footprint corresponding with human development index: eco-sustainable human development index, *International Journal of Sustainable Development & World Ecology*, 20, 9–19, 2013.
- Veraverbeke, S., Rogers, B. M., Goulden, M. L., Jandt, R. R., Miller, C. E., Wiggins, E. B., and Randerson, J. T.: Lightning as a major driver of recent large fire years in North American boreal forests, *Nature Climate Change*, 7, 529–534, 2017.
- Verjans, V., Franzke, C. L., Lee, S.-S., Kim, I.-W., Tilmes, S., Lawrence, D. M., Vitt, F., and Li, F.: Quantifying CO2 forcing effects on lightning, wildfires, and climate interactions, *Science Advances*, 11, eadt5088, 2025.
- 1075 Viovy, N.: CRUNCEP version 7-atmospheric forcing data for the community land model, Research Data Archive at the National Center for Atmospheric Research, Computational and Information Systems Laboratory, 10, 2018.
- Vitolo, C., Di Giuseppe, F., Barnard, C., Coughlan, R., San-Miguel-Ayanz, J., Libertá, G., and Krzeminski, B.: ERA5-based global meteorological wildfire danger maps, *Scientific data*, 7, 216, 2020.
- 1080 Williams, A. P., Abatzoglou, J. T., Gershunov, A., Guzman-Morales, J., Bishop, D. A., Balch, J. K., and Lettenmaier, D. P.: Observed impacts of anthropogenic climate change on wildfire in California, *Earth’s Future*, 7, 892–910, 2019.
- Williams, K., Copesey, D., Blockley, E., Bodas-Salcedo, A., Calvert, D., Comer, R., Davis, P., Graham, T., Hewitt, H., Hill, R., et al.: The Met Office global coupled model 3.0 and 3.1 (GC3. 0 and GC3. 1) configurations, *Journal of Advances in Modeling Earth Systems*, 10, 357–380, 2018.
- 1085 Yue, C., Ciaia, P., Cadule, P., Thonicke, K., Archibald, S., Poulter, B., Hao, W., Hantson, S., Mouillot, F., Friedlingstein, P., et al.: Modelling the role of fires in the terrestrial carbon balance by incorporating SPITFIRE into the global vegetation model ORCHIDEE–Part 1: simulating historical global burned area and fire regimes, *Geoscientific Model Development*, 7, 2747–2767, 2014.
- Zou, Y., Wang, Y., Ke, Z., Tian, H., Yang, J., and Liu, Y.: Development of a REgion-specific ecosystem feedback fire (RESFire) model in the Community Earth System Model, *Journal of Advances in Modeling Earth Systems*, 11, 417–445, 2019.

1090 Zubkova, M., Boschetti, L., Abatzoglou, J. T., and Giglio, L.: Changes in fire activity in Africa from 2002 to 2016 and their potential drivers, *Geophysical research letters*, 46, 7643–7653, 2019.