

An extension of the WeatherBench 2 to binary hydroclimatic forecasts

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Abstract: Binary forecasts on hydroclimatic extremes play a critical part in disaster prevention and risk management. While the recent WeatherBench 2 provides a versatile framework for the verification of deterministic and ensemble forecasts, this paper presents an extension to binary forecasts on the occurrence versus non-occurrence of hydroclimatic extremes. Specifically, ~~sixteen~~ **seventeen** verification metrics on the accuracy and discrimination of binary forecasts are employed and scorecards are generated to showcase the predictive performance. A case study is devised for binary forecasts of wet and warm extremes obtained from both deterministic and ensemble forecasts generated by three data-driven models, i.e., Pangu-Weather, GraphCast and FuXi, and two numerical weather prediction products, i.e., ECMWF's IFS HRES and IFS ENS. The results show that the receiver operating characteristic skill score (ROCSS) serves as a suitable metric due to its relative insensitivity to the rarity of hydroclimatic extremes. For wet extremes, the GraphCast tends to outperform the IFS HRES with the total precipitation of ERA5 **reanalysis** data as **the** ground truth. For warm extremes, the Pangu-Weather, GraphCast and FuXi tends to be more skilful than the IFS HRES within 3-day lead time but become less skilful as lead time increases. In the meantime, the IFS ENS tends to provide skilful forecasts of both wet and warm extremes at different lead times and at the global scale. Through diagnostic plots of forecast time series at selected grid cells, it is observed that at longer lead times, forecasts generated by data-driven models tend to be smoother and less skilful compared to those generated by physical models. Overall, the extension of the WeatherBench 2 facilitates more comprehensive comparisons of hydroclimatic forecasts and provides useful information for forecast applications.

Keywords: Binary forecast; forecast verification; warm extreme; wet extreme; forecast skill; scorecard.

25 1 Introduction

Accurate numerical weather prediction (NWP) is of great importance to the economy and society (Bi et al., 2023; Lam et al., 2023; Bauer et al., 2015). Conventionally, physical NWP models formulate the governing equations of coupled physical processes in land, ocean and atmosphere and therefore predict weather conditions in the near future based on predetermined initial meteorological fields (Lam et al., 2023; Bauer et al., 2015). Due to advances in remote sensing, data assimilation and
30 computational infrastructure, physical NWP models have witnessed steady improvements and been extensively employed in operational forecasting (Bauer et al., 2020). For example, the European Centre for Medium-range Weather Forecast (ECMWF) operates the Integrated Forecast System (IFS) that has implemented a ~~significant~~ [remarkable](#) resolution upgrade and methodology for high-resolution forecasts (HRES) and ensemble forecasts (ENS) at the horizontal resolution of 0.1 degrees since January 2016 (Balsamo et al., 2023).

35 Data-driven NWP models have recently gained increasing popularity in hydroclimatic forecasting (Ben Bouallège et al., 2024; Rasp et al., 2024; de Burgh-Day and Leeuwenburg, 2023; Xu et al., 2024a). Early models, such as the UNet architecture-based cubed sphere projection (Weyn et al., 2020) and deep Resnet architecture-based models (Clare et al., 2021; Rasp and Thuerey, 2021), were of moderate spatial-temporal resolution and forecast skill. Recent deep learning models, such as graph neural network (Keisler, 2022) and FourCastNet (Pathak et al., 2022), began to match operational NWP models in resolution
40 and skills. Pangu-Weather (Bi et al., 2023) and GraphCast (Lam et al., 2023) even outperformed the HRES in terms of some deterministic metrics. The Neural General Circulation Models (NeuralGCM) that integrates data-drive and physical modules is considered to be the first hybrid model obtaining competitive or better scores than the HERS (Kochkov et al., 2024). The GenCast generates global ensemble forecasts that are comparative or even more skilful than the ENS (Price et al., 2025).

There is a growing demand to verify the capability of physical and data-driven models in generating skilful hydroclimatic
45 forecasts (Olivetti and Messori, 2024a; Zhong et al., 2024; Ben Bouallège et al., 2024). In response to the need of a unified benchmark, the WeatherBench has been established to host a common dataset of forecasts and observations and utilizes popular evaluation metrics for forecast comparisons (Rasp et al., 2020). Owing to rapid advances in data-driven NWP models, the WeatherBench 2 has been developed to support global medium-range forecast verification (Rasp et al., 2024). By following
50 established practices in the World Meteorological Organisation (WMO), the WeatherBench 2 pays attention to both deterministic and ensemble forecasts generated by physical and data-driven NWP models (Jin et al., 2024). Forecast verification is performed by an open-source Python code and publicly available, cloud-optimized ground-truth and baseline datasets (Jin et al., 2024; Olivetti and Messori, 2024b; Rasp et al., 2024).

Besides deterministic and ensemble forecasts, there is a demand of binary forecasts in disaster prevention and risk management (Ben Bouallège et al., 2024; Larraondo et al., 2020). ~~The importance of binary forecasts arises from the~~

operational need for predicting the occurrence versus non-occurrence of hydroclimatic extremes (Jolliffe and Stephenson, 2012; Larraondo et al., 2020; Rasp et al., 2020). For operational applications such as disaster warning, the emphasis is usually on usually pay attention to the occurrence versus non-occurrence of certain hydroclimatic extremes instead of their precise magnitude (Larraondo et al., 2020; Rasp et al., 2020). Binary forecasts match this need and can be directly used for decision-making (Jolliffe and Stephenson, 2012; Larraondo et al., 2020). In the meantime, binary forecasts meet this demand by emphasizing the ability to capture hydroclimatic extremes, ensuring that models are not rewarded for merely minimizing average errors and unrealistically smooth forecasts (Ferro and Stephenson, 2011; Rasp et al., 2020). Therefore, this paper aims to extend the WeatherBench 2 to binary forecasts. The objectives are: 1) to account for verification metrics on binary forecasts derived from global precipitation and temperature forecasts; 2) to present scorecards to showcase the predictive performance on wet and warm extremes; and 3) to examine the sensitivity of different metrics to predefined thresholds of hydroclimatic extremes. As will be shown in the methods and results, the extension facilitates an effective intercomparison among binary forecasts of hydroclimatic extremes generated by both data-driven and physically-based physical models.

2 Forecasts and metrics in the WeatherBench 2

2.1 Forecast datasets

The WeatherBench 2 presents a benchmark for verifying and comparing the performance of data-driven and physical NWP models (Rasp et al., 2024). On its website (<https://weatherbench2.readthedocs.io>), there is a database containing past forecasts in the year 2020:

1) The HRES generated by the ECMWF's IFS is widely regarded as one of the best global deterministic weather forecasts (Rasp et al., 2024). It offers 10-day forecasts at the horizontal resolution of 0.1 degrees with 137 vertical levels (Balsamo et al., 2023). In the WeatherBench_2, the HRES is primarily used as the main-baseline for comparing the performance of data-driven models.

2) The ENS generated by the IFS's ensemble version is widely known as one of the best global ensemble weather forecasts. It consists of 1 control member and 50 perturbed members (Balsamo et al., 2023). In the WeatherBench_2, the ENS is also used-serves as an important baseline-, with The-the average-mean value of over the 50 members (i.e., ENS Mean)-, being extensively usedis-considered-as-an-additional-baseline (Rasp et al., 2024).

3) The ERA5 forecasts are hindcasts generated by the exact IFS that is employed to create the renowned ERA5 of historical global climate conditions (Hersbach et al., 2020). They consist of 10-day hindcasts at the horizontal resolution of 0.25 degrees.

64) The two sets of 10-day forecasts generated by the GraphCast includes 6 upper-air variables at a maximum of 37 vertical levels and 5 surface variables at the horizontal resolution of 0.25 degrees (Lam et al., 2023). The GraphCast is based on the architecture of graph neural network. It runs autoregressively to forecast atmospheric states for the next time step based on states from the previous two time steps at the temporal resolution of 6 hours. Similarly, there are two sets of GraphCast forecasts generated from the ERA5 and HRES initializations (Rasp et al., 2024).

7) The 10-day global forecasts generated by the spherical convolutional neural networks (CNNs) are composed of 6 upper-air variables at 13 vertical levels (Esteves et al., 2023). The Spherical CNNs extend the CNNs to the sphere domains through leveraging the spherical convolutions as the primary linear operation. It produces forecasts at the longitude resolution of 1.4 degrees and the latitude resolution of 0.7 degrees.

9) The 15-day deterministic and ensemble forecasts generated by NeuralGCM are composed of 7 upper-air variables at 37 vertical levels at the temporal resolution of 12 hours (Kochkov et al., 2024). The NeuralGCM integrates the differential dynamical core and the learned physics module. The deterministic version is trained in the horizontal resolution of 0.7 degrees. The ensemble version is trained for the horizontal resolution of 1.4 degrees and is run to produce 50 members.

2.2 Verification metrics

The WeatherBench 2 takes into consideration in total 6 metrics for deterministic forecasts and 6 metrics for ensemble forecasts, as shown in Table 1. In forecast verification, the ERA5 reanalysis data is used as the ground truth for verifying the data-driven models. For the sake of fair comparison with the data-driven models, the initial conditions of the IFS HRES is used as the ground truth for verifying the verification of IFS forecasts (Lam et al., 2023). As precipitation is not available for the IFS HRES's initial conditions, the total precipitation of ERA5 reanalysis data is used as the ground truth data for all models. In the initial version of WeatherBench 2, the verification is conducted for forecasts initialized at 00 and 12 UTC for the period from 1 January 2020 to 31 December 2020. All forecasts, baseline data and ground truth data are resampled to the horizontal resolution of 1.5 degrees that is used as the standard resolution for forecasts verification by the WMO and ECWMF (Rasp et al., 2024).

Table 1. Verification Metrics metrics for deterministic and ensemble forecasts in the WeatherBench 2.

| Forecast | Metric | [min, max] | Optimal value |
|---------------|---|----------------------|---------------|
| Deterministic | Root mean square error (RMSE) | $[0, +\infty)$ | 0 |
| | Mean square error | $[0, +\infty)$ | 0 |
| | Mean absolute error | $[0, +\infty)$ | 0 |
| | Bias | $(-\infty, +\infty)$ | 0 |
| | Anomaly correlation coefficient | $[-1, 1]$ | 1 |
| | Stable Equitable Error in Probability Space (SEEPS) | $[0, 1]$ | 0 |
| Ensemble | Continuous ranked probability score (CRPS) | $[0, +\infty)$ | 0 |

| | | |
|--------------------------------|----------------|---|
| Ranked probability score (RPS) | $[0, +\infty)$ | 0 |
| Spread-Skill Ratio | $[0, 1]$ | 1 |
| Energy score | $[0, +\infty)$ | 0 |
| Brier score (BS) | $[0, 1]$ | 0 |
| Ignorance score | $[0, +\infty)$ | 0 |

3 Verification of binary hydroclimatic forecasts

3.1 Conversion to binary forecasts

Binary forecasts on the occurrence versus non-occurrence of target events can be generated from deterministic and ensemble forecasts by using predefined thresholds of hydroclimatic events (Ben Bouallègue et al., 2024). In operational applications, binary forecasts of extreme precipitation events as the critical variables to understand and forecast hydroclimatic processes such as floods and heatwaves, can respectively be derived from the precipitation and temperature are of concern in operational forecasts (Huang and Zhao, 2022; Lang et al., 2014; Zhao et al., 2022; Slater et al., 2023). As to precipitation, the 90th percentile of the 24-hour accumulation of total precipitation (TP24h) is considered as the threshold, above which the TP24h is considered as the wet extreme (North et al., 2013). As to temperature, the 90th percentile of the 24-hour maximum of 2m temperature (T2M24h) is set as the threshold, above which the T2M24h is categorized as the warm extreme (Olivetti and Messori, 2024b). It is noted that the thresholds at each grid cell are separately calculated separately for each grid cell to obtain an equal number of extreme samples. Given the pre-defined threshold (q), deterministic forecasts are converted into either 0 or 1:

$$I(f_n > q) = \begin{cases} 1, & f_n > q \\ 0, & \text{otherwise} \end{cases} \quad (1)$$

where f_n represents the n -th deterministic forecast. By contrast, In the meantime, ensemble forecasts are converted into forecast probabilities by using the Weibull's plotting position (Makkonen, 2006):

$$p_{f_n} = \frac{\sum_{m=1}^M I(f_{n,m} > q)}{M + 1} \quad (2)$$

where $f_{n,m}$ is the m -th member of the n -th ensemble forecasts and M is the number of ensemble members.

In relation to the corresponding observations, The comparison of binary forecasts against the corresponding observations facilitate can be divided into four categories, i.e., hits true positives (a), false positivesalarms (b), false negativesmisses (c) and true negativescorrect rejections (d), as shown in Table 2 (Larraondo et al., 2020). Specifically, The the true positives hits representindicate that target occurrences events that are successfully forecasted; the false positivesalarms are-indicate non-occurrences non-events that are incorrectly forecasted as occurrencesevents; the false negatives misses denoteindicate target occurrences-events that are incorrectly forecasted as non-occurrencesnon-events; and the true negativescorrect rejections

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indicate represent non-occurrences that non-events that are correctly forecasted as non-occurrences non-events. The proportion of the observed events to the total number of events and non-events is the base rate $((a + c)/N)$, with lower values often corresponding to events that are more extreme (Ferro and Stephenson, 2011).

Table 2. Contingency table for binary forecasts.

| Observed occurrences events | | Observed non-occurrences events | | Total |
|-----------------------------------|---|-----------------------------------|---|---------------------|
| Forecasted occurrences events | $a = \begin{cases} \sum_{n=1}^N I(f_n > q p_{o_n} = 1), & \text{if } M = 1 \\ \sum_{n=1}^N I(p_{f_n} > p_e p_{o_n} = 1), & \text{if } M \geq 2 \end{cases}$ | Forecasted non-occurrences events | $b = \begin{cases} \sum_{n=1}^N I(f_n > q p_{o_n} = 0), & \text{if } M = 1 \\ \sum_{n=1}^N I(p_{f_n} > p_e p_{o_n} = 0), & \text{if } M \geq 2 \end{cases}$ | $a + b$ |
| Forecasted non-occurrences events | $c = \begin{cases} \sum_{n=1}^N I(f_n \leq q p_{o_n} = 1), & \text{if } M = 1 \\ \sum_{n=1}^N I(p_{f_n} \leq p_e p_{o_n} = 1), & \text{if } M \geq 2 \end{cases}$ | Forecasted occurrences events | $d = \begin{cases} \sum_{n=1}^N I(f_n \leq q p_{o_n} = 0), & \text{if } M = 1 \\ \sum_{n=1}^N I(p_{f_n} \leq p_e p_{o_n} = 0), & \text{if } M \geq 2 \end{cases}$ | $c + d$ |
| Total | $a + c$ | Total | $b + d$ | $a + b + c + d = N$ |

Where $M = 1$ and $M \geq 2$ respectively denote represent the deterministic forecasts and ensemble forecasts; N is the number of pairs of observations and forecasts for verification; p_{o_n} represents the binary observation with which is a value of 1 for the occurrences event and 0 for non-occurrences event; p_e denotes the probability thresholds above which occurrences the events are forecasted to occur for ensemble forecasts.

3.2 Verification metrics for binary forecasts

Given the challenges posed by varying hydroclimatic extremes and imbalanced samples, in total 4617 metrics are utilized to examine the performance of binary forecasts (Jolliffe and Stephenson, 2012; North et al., 2013). Notably, There there are 7 8 base-rate-dependent metrics and 9 base-rate-independent metrics. On the one hand, The the base-rate-dependent metrics provide facilitate insights into the performance in relation to varying frequency of extreme events (Jolliffe and Stephenson, 2012). On the other hand, the base-rate-independent metrics are suitable applicable to for comparing forecasts across different climate regions or time periods, in which the frequency of extreme events differs substantially (Ferro and Stephenson, 2011; Jacox et al., 2022). Their equations, ranges and optimal values are presented in Table 3.

Table 3. Metrics for binary forecasts.

| Metric | Equation | [min, max] | Optimal value | Reference |
|-----------------------------|----------|------------|---------------|-----------|
| Base-rate-dependent metrics | | | | |

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| Accuracy (ACC), proportion correct | $ACC = \frac{a + d}{N}$ | [0, 1] | 1 | (Finley, 1884) |
| Success ratio (SR), precision | $SR = \frac{a}{a + b}$ | [0, 1] | 1 | (Lagadec et al., 2016) |
| Critical success index (CSI), threat score, Gilbert score | $CSI = \frac{a}{a + b + c}$ | [0, 1] | 1 | (Donaldson et al., 1975; Gilbert, 1884) |
| Heidke skill score (HSS), Cohen's Kappa | $HSS = \frac{a + d - a_p - d_p}{N - a_p - d_p}, d_p = \frac{(b + d)(c + d)}{N}$ | [-1, 1] | 1 | (Gomis-Cebolla et al., 2023; Heidke, 1926) |
| Gilbert skill score (GSS), equitable threat score | $GSS = \frac{a - a_r}{a + b + c - a_r}, a_r = \frac{(a + b)(a + c)}{N}$ | [-1/3, 1] | 1 | (Gilbert, 1884; Schaefer, 1990) |
| Heidke skill score (HSS), Cohen's Kappa | $HSS = \frac{a + d - a_r - d_r}{N - a_r - d_r}, d_r = \frac{(b + d)(c + d)}{N}$ | [-1, 1] | 1 | (Gomis-Cebolla et al., 2023; Heidke, 1926) |
| Extreme dependence score (EDS) | $EDS = \frac{\ln[(a + c)/N]p - \ln H}{\ln[(a + c)/N]p + \ln H}, p = \frac{a + c}{N}$ | [-1, 1] | 1 | (Primo and Ghelli, 2009; Stephenson et al., 2008) |
| Symmetric extreme dependence score (SEDS) | $SEDS = \frac{\ln[(a + b)/N]q - \ln H}{\ln[(a + c)/N]p + \ln H}, q = \frac{a + b}{N}$ | [-1, 1] | 1 | (Orozco López et al., 2010) |
| Potential relative economic value (REV) | $REV = \max_{0 \leq p \leq 1} \frac{\min\{a + c, r\} - [(a + b)r + c]}{\min\{a + c, r\} - (a + c)r}$ | [0, 1] | 1 | (Richardson, 2006, 2000; Wilks, 2001) |
| Base-rate-independent metrics | | | | |
| Hit rate (H), sensitivity, recall, probability of detection | $H = \frac{a}{a + c}$ | [0, 1] | 1 | (Swets, 1986) |
| False alarm rate (F), probability of false detection | $F = \frac{b}{b + d}$ | [0, 1] | 0 | (Donaldson et al., 1975) |
| Specificity, true negative rate (TNR) | $TNR = \frac{d}{b + d}$ | [0, 1] | 1 | (Agrawal et al., 2023) |
| Odds ratio skill score (ORSS), Yule's Q | $ORSS = \frac{ad - bc}{ad + bc}$ | [-1, 1] | 1 | (Stephenson, 2000) |
| Peirce's skill score (PSS), Hanssen and Kuipers discriminant | $PSS = \frac{ad - bc}{(a + c)(b + d)} = H - F$ | [-1, 1] | 1 | (Peirce, 1884) |
| Extremal dependence index (EDI) | $EDI = \frac{\ln F - \ln H}{\ln F + \ln H}$ | [-1, 1] | 1 | (Ferro and Stephenson, 2011) |
| Symmetric extremal dependence index (SEDI) | $SEDI = \frac{\ln F - \ln H + \ln(1 - H) - \ln(1 - F)}{\ln F + \ln H + \ln(1 - H) + \ln(1 - F)}$ | [-1, 1] | 1 | (Ferro and Stephenson, 2011) |
| Area under receiver operating characteristic (ROC) curve (AUC) | $AUC = \int_0^1 HdF \phi\left(\frac{\Phi^{-1}(H) - \Phi^{-1}(F)}{\sqrt{2}}\right)$ | [0, 1] | 1 | (Swets, 1986) |
| ROC skill score (ROCSS) | $ROCSS = 2(AUC - 0.5)$ | [-1, 1] | 1 | (Swets and Swets, 1986) |

Where a , b , c and d respectively denote the number of true positives, false positives, false negatives and true negatives, with the equations shown in Table 2; N is the number of pairs of observations and forecasts; p denotes the probability thresholds above which the events are forecasted to occur for ensemble forecasts; r represents the cost-loss ratio for calculating the relative economic value; all calculation equations of other variables can be found in this table.

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The 7-8 base-rate-dependent metrics in Table 3 are influenced by the underlying distribution of observed events and non-events (Jolliffe and Stephenson, 2012). The accuracy is calculated as the ratio between the number of ~~hits-true positives~~ and the total number of events and non-events (Finley, 1884). The success ratio (SR) measures the number of ~~true positives hits~~ divided by the number of forecasted events (Lagadec et al., 2016). The critical success index (CSI) is the number of ~~true positives hits~~ divided by the total number of forecasted and observed events (Chakraborty et al., 2023; Gilbert, 1884; Donaldson et al., 1975). ~~The Gillert skill score (GSS) evaluates the fraction of true positives over the observed and forecasted events after adjusting for the random true positives (Chen et al., 2018; Coelho et al., 2022).~~ The Heidke skill score (HSS) measures the accuracy relative to that of the random forecasts (Gomis-Cebolla et al., 2023). ~~The Gillert skill score (GSS) evaluates the correctly predicted fraction of the observed and forecasted events after adjusting for the random hits (Chen et al., 2018; Coelho et al., 2022). Converging to a meaningful limit,~~ The extreme dependency score (EDS) (Stephenson et al., 2008) and the symmetric extreme dependency score (SEDS) (Orozco López et al., 2010) ~~are suitable can measure the general performance of for~~ binary forecasts ~~of for~~ rare events. ~~The potential relative economic value (REV) quantifies the potential value of a forecast over a range of different probability thresholds (p) to make decision (Richardson, 2006, 2000; Wilks, 2001).~~ It compares the saved expense using the forecasts instead of climatology relative to the saved expense using the perfect forecast (Price et al., 2025).

The 9 base-rate-independent metrics in Table 3 are valuable for rare events due to their stability to the variation in the proportion of observed events (Ferro and Stephenson, 2011). The hit rate and false alarm rate respectively quantify the proportion of ~~true positiveshits~~ in observed events and the proportion of false ~~alarms-positives~~ in observed non-events (Swets, 1986). The specificity measures the percentage of ~~true negatives correct rejections~~ to observed non-events (Agrawal et al., 2023). The odds ratio skill score (ORSS) examines the improvement over the random forecasts, ~~emphasizing the balance between positive and negative samples~~ (Stephenson, 2000). The Peirce's skill score (PSS) has similar formulation to HSS but does not depend on event frequency (Chakraborty et al., 2023). ~~For deterministic forecasts, the PSS equals to the maximum value of REV when the cost-loss ratio equals to the base rate (Richardson, 2006).~~ The extremal dependence index (EDI) and the symmetric extremal dependence index (SEDI) are designed to be nondegenerate ~~to measure the predictive performance for rare events.~~ (Ferro and Stephenson, 2011). The receiver operating characteristic (ROC) examines the discrimination between ~~true positiveshits~~ and false ~~positives-alarms~~, quantified by the area under the ROC curve (AUC) (Swets, 1986). ~~The ROC skill score (ROCSS) compares the discriminative ability over random forecasts. A higher ROC skill score (ROCSS) indicates better predictive skill. For probabilistic forecasts, the ROCSS can be calculated by considering the hit rate and false alarm rate for all possible thresholds of probability (Huang and Zhao, 2022).~~

Among the ~~seventeen~~ 17 metrics, the ROCSS is base-rate-independent and suitable ~~simultaneously~~ for both deterministic

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and probabilistic forecasts of binary events. By contrast, the other metrics need some predefined probability thresholds to convert the probabilistic forecasts into deterministic forecasts. Therefore, the ROCSS is selected as the main primary verification metric in the analysis. For probabilistic forecasts, the ROCSS can be calculated by considering the hit rate and false alarm rate for all possible thresholds of probability (Huang and Zhao, 2022). It is noted that A higher ROCSS values indicates better predictive forecast skill.

3.3 Forecast verification

Considering data availability and forecast settings, the verification focuses on attention is paid to 8 sets of forecasts: IFS’s HRES, ENS and ENS Mean; operational forecasts from Pangu-Weather, GraphCast; and hindcasts from Pangu-Weather, GraphCast and FuXi, i.e., IFS HRES, IFS ENS, IFS ENS Mean, Pangu-Weather (operational), GraphCast (operational), Pangu-Weather, GraphCast and FuXi (Rasp et al., 2024). The ground truth, spatial resolution, initial forecast time and verification period are selected by following the WeatherBench 2. A set of predefined thresholds ranging from the 80th to 99th percentiles of the ground truth data in 2020 are considered for sensitivity analysis (Olivetti and Messori, 2024b; North et al., 2013). For the comparison at individual grid cells, the 17 metrics are one by one calculated. Furthermore, the 17 metrics are calculated using the area-weighting method for the regions pre-determined by the ECMWF’s scorecards (Rasp et al., 2024). For comparison at the grid scale, the 17 metrics are computed separately for each grid cell. To facilitate comparisons at regional to global scales, the 1617 metrics are calculated based on using the area-weighting method over grid cells based on the scores that are separately calculated for each grid cell (Rasp et al., 2024). The regions are defined by the ECMWF’s scorecards, as illustrated in Table 4.

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Table 4. Regions that are included in the ECMWF’s scorecards.

| Region | Range | Region | Range |
|-------------------------------------|--|----------------|--|
| Northern hemisphere (extra-tropics) | latitude $\geq 20^{\circ}$ | Europe | $35^{\circ} \leq \text{latitude} \leq 75^{\circ}$, $-12.5^{\circ} \leq \text{longitude} \leq 42.5^{\circ}$ |
| Southern hemisphere (extra-tropics) | latitude $\leq -20^{\circ}$ | North America | $25^{\circ} \leq \text{latitude} \leq 60^{\circ}$, $-120^{\circ} \leq \text{longitude} \leq -75^{\circ}$ |
| Tropics | $-20^{\circ} \leq \text{latitude} \leq 20^{\circ}$ | North Atlantic | $25^{\circ} \leq \text{latitude} \leq 60^{\circ}$, $-70^{\circ} \leq \text{longitude} \leq -20^{\circ}$ |
| Extra-tropics | $ \text{latitude} \geq 20^{\circ}$ | North Pacific | $25^{\circ} \leq \text{latitude} \leq 60^{\circ}$, $145^{\circ} \leq \text{longitude} \leq -130^{\circ}$ |
| Arctic | latitude $\geq 60^{\circ}$ | East Asia | $25^{\circ} \leq \text{latitude} \leq 60^{\circ}$, $102.5^{\circ} \leq \text{longitude} \leq 150^{\circ}$ |
| Antarctic | latitude $\leq -60^{\circ}$ | AusNZ | $-45^{\circ} \leq \text{latitude} \leq -12.5^{\circ}$, $120^{\circ} \leq \text{longitude} \leq 175^{\circ}$ |

AusNZ: Australia and New Zealand.

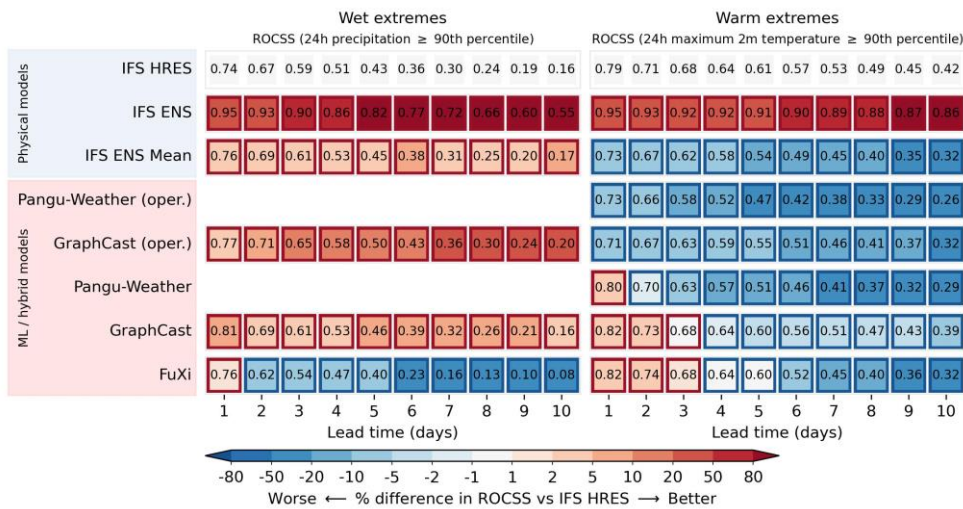


Figure 1. Globally area-weighted ROCSS for wet and warm extremes. The oper. denotes the operational version. The red and blue borders indicate significantly different performances compared to the IFS HRES at the significance level of 0.05.

Scorecards of the area-weighted ROCSS for wet extremes relative to the IFS HRES baseline are illustrated by region in Figure 2. Overall, the IFS ENS stands out across different regions and lead times. The GraphCast (operational) tends to outperform the IFS HRES. The GraphCast tend to be better than the IFS HRES in Southern Hemisphere (extra-tropics), Arctic, Antarctic, Europe, North Pacific, East Asia and AusNZ. In Europe, at the lead times of 3 and 10 days, the ROCSS is respectively 0.73 and 0.19 for the IFS HRES, 0.96 and 0.64 for the IFS ENS, 0.76 and 0.23 for the GraphCast (operational), 0.77 and 0.22 for the GraphCast and 0.69 and 0.11 for the FuXi. In the meantime, the FuXi tends to outperform the IFS HRES in the Southern Hemisphere (extra-tropics), tropics, North Atlantic and AusNZ at lead time less than 3 days. Except for the Arctic and Antarctic, the IFS ENS Mean tends to be better than the IFS HRES. The GraphCast (operational) is comparable to the IFS ENS Mean and marginally better in the polar regions. In the Antarctic region, the ROCSS is 0.63 and 0.06 for the IFS HRES, 0.59 and 0.01 for the IFS ENS Mean and 0.66 and 0.06 for the GraphCast (operational) at lead times of 3 and 10 days.

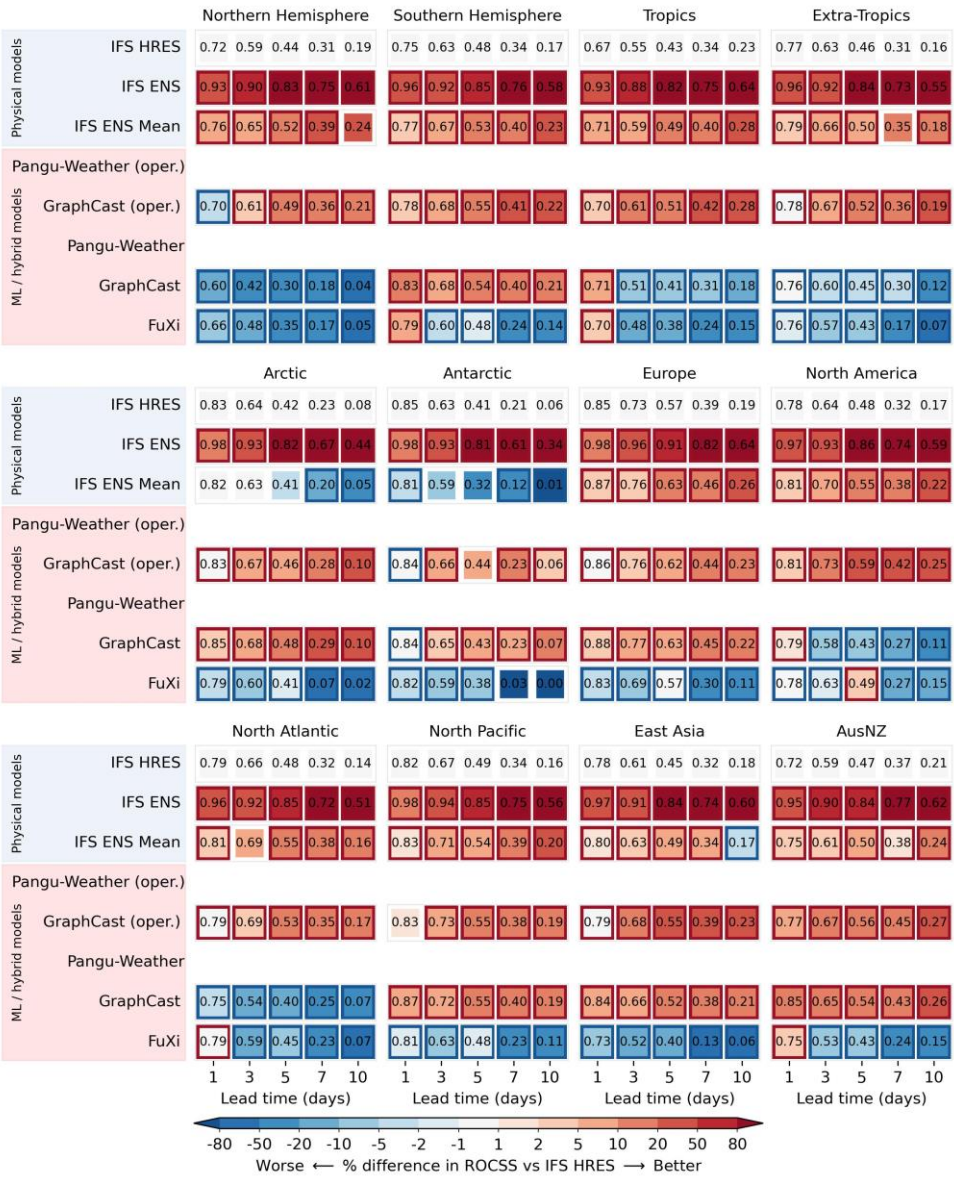


Figure 2. Regionally area-weighted ROCSS of different forecasts for wet extreme. The red and blue borders indicate significantly different performance compared to the IFS HRES at the significance level of 0.05.

Scorecards of the regionally area-weighted ROCSS for warm extremes relative to the IFS HRES baseline are showcased in Figure 3. The Pangu-Weather, GraphCast and FuXi tend to outperform the IFS HRES within 3-day lead time except for the

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Arctic and Antarctic. These results are consistent with the results of a previous study on forecast accuracy of [the magnitude for](#) warm extremes (Olivetti and Messori, 2024b). In the North America, North Atlantic, North Pacific, East Asia and AusNZ, the GraphCast and FuXi tend to outperform the IFS HRES at longer lead times even up to 10 days. The ROCSS in the North Atlantic is respectively 0.39, 0.58 and 0.49 for the IFS HRES, GraphCast and FuXi at the 10-day lead time. On the other hand, the performances of all data-driven forecasts tend to be worse than that of the IFS HRES in the Arctic and Antarctic. In Europe, the ROCSS is respectively 0.78, 0.71, 0.76 and 0.75 for the IFS HRES, Pangu-Weather, GraphCast and FuXi at 5-day lead time. As averaging the ensemble members can filter unpredictable features to get smoother forecasts, it is not surprising that the IFS ENS Mean does not always perform as well as the IFS HRES and IFS ENS for warm extremes (Ben Bouall gue et al., 2024).

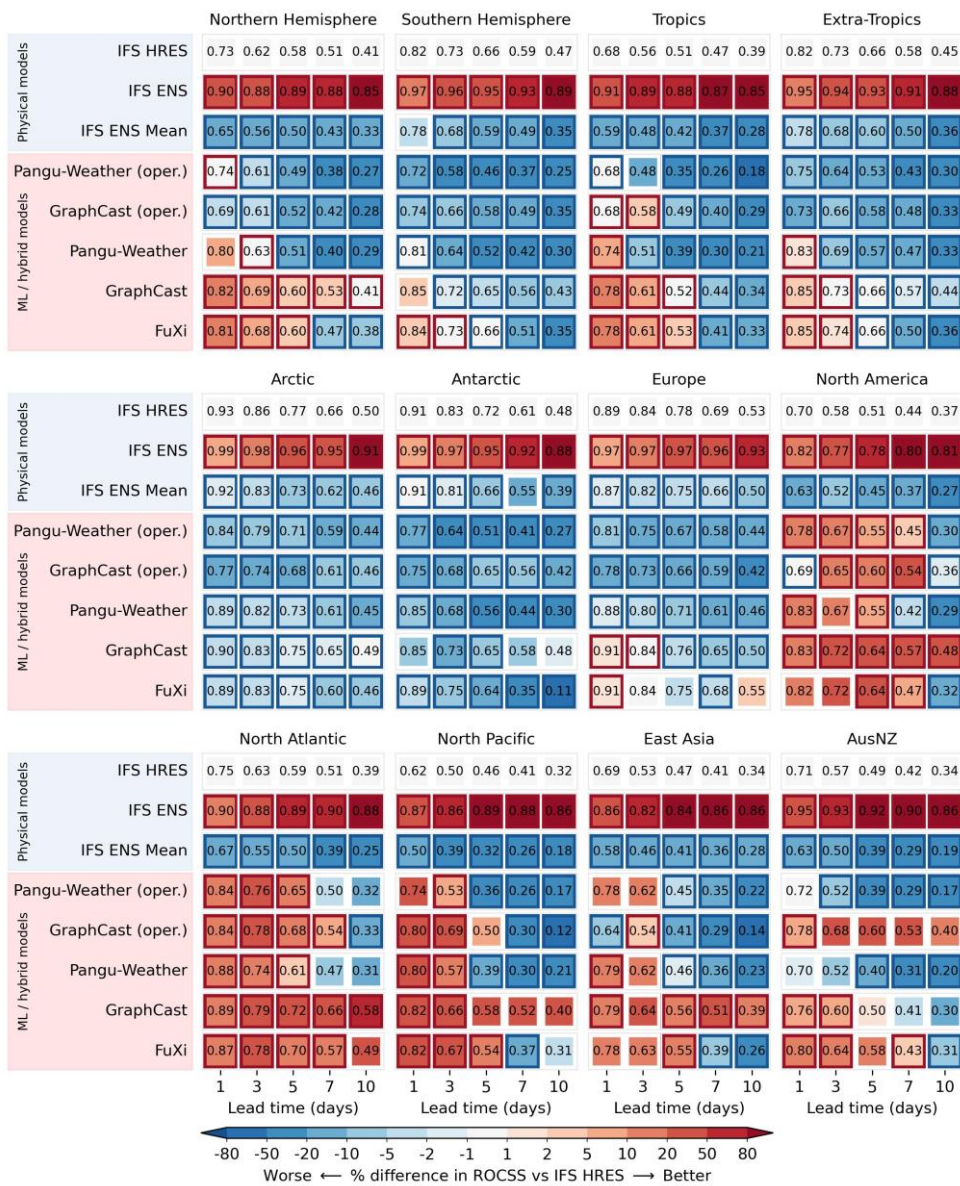


Figure 3. As for Figure 2, but for warm extremes.

4.2 Predictive performance of wet extremes

The differences in the ROCSS for wet extremes in comparison with the IFS HRES baseline are illustrated in Figure 4. Overall, the IFS ENS tends to outperform the IFS HRES at most grid cells across the globe. Except for the Northern Africa

and Arabian Peninsula, the GraphCast’s operational forecasts are comparable or more skilful than the IFS HRES. The GraphCast is not as skilful as the IFS HRES in more grid cells, such as [the Northern Africa](#), Central Australia and Central Asia. The FuXi tends to be less skilful than the IFS HRES in most grid cells, such as the [Northern Africa](#), Atlantic and Pacific. As the lead time increases, the IFS ENS and GraphCast (operational) are observed to outperform the IFS HRES, while the GraphCast and FuXi underperform. These results are consistent with the results of Figure 1 and Figure 2. In Northern Africa, forecasts of the three data-driven models tend to be less skilful than the IFS HRES and IFS ENS. As the GraphCast and FuXi exhibit no hits ~~or~~ [and so many false alarms-positives](#) for many of or even almost all the grid cells in this region, the ROCSS is nearly ~~zero~~ [-1](#) so that their forecasts tend to be worse than the IFS HRES in the Northern Hemisphere (extra-tropics) and Tropics.

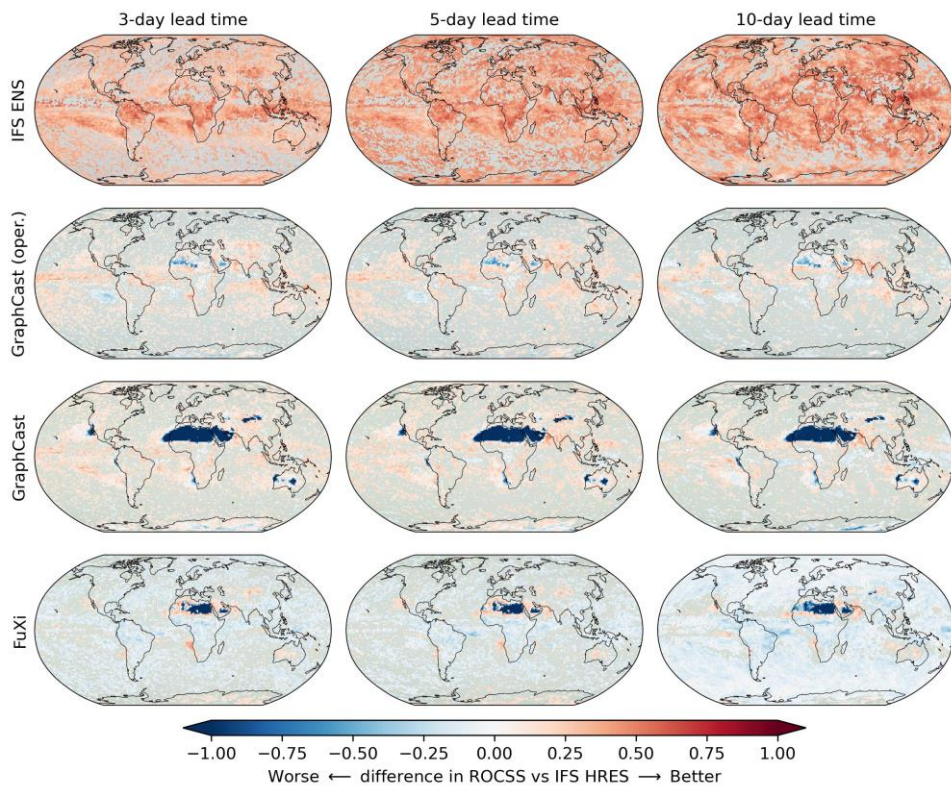


Figure 4. Differences of IFS ENS, GraphCast (operational), GraphCast and FuXi in ROCSS to the IFS HRES for wet extremes at each grid cell. The grey colour indicates grid with no statistically significant differences at the significance level of 0.1.

The time series for 24-hour accumulation of total precipitation from different forecasts initialized at 00 UTC are shown for three ~~selected~~ grid cells in Figure 5. The grid cells A, B and C are selected respectively due to the better, close and worse performance of data-driven models in relative to the IFS HRES. Overall, data-driven models can capture the temporal dynamics of precipitation but their forecasts are smoother than the IFS HRES (Zhong et al., 2024; Xu et al., 2024b). As lead time increases, the skill of IFS ENS and IFS ENS Mean tend to decrease more slowly than that of IFS HRES, while the skill of GraphCast and FuXi reduces more rapidly. For the three grid cells, the five sets of forecasts have close number of correct negatives; the IFS HRES and GraphCast show more hits; the IFS HRES are more capable of capturing the wet extremes but tends to produce more false alarms; the IFS ENS Mean and FuXi tend to underestimate the wet extremes, resulting in more misses and fewer false alarms. For grid cell A, at the lead times of 3 and 10 days, the ROCSS is respectively 0.63 and 0.34 for the IFS HRES, 0.92 and 0.76 for the IFS ENS, 0.62 and 0.43 for the IFS ENS Mean, 0.75 and 0.38 for the GraphCast and 0.48 and 0.16 for the FuXi. For grid cells B and C, the numbers of hits and false alarms of FuXi are zero at 10-day lead time, leading to zero values of ROCSS. For grid cells A and B, the five sets of forecasts have nearly equal number of true negatives; the IFS HRES show more true positives but more false negatives; the GraphCast is more capable of capturing the wet extremes but tends to produce more false positives; the IFS ENS Mean and FuXi tend to underestimate the wet extremes, resulting in more false negatives but fewer false positives. For grid cell C that is located in the Northern Africa, the GraphCast and FuXi tend to overestimate the low precipitation and underestimate the high precipitation, leading to zero numbers of true negatives for the FuXi and zero numbers of false negatives for both. At the lead times of 3 and 10 days, the ROCSS is respectively 0.48 and 0.09 for the IFS HRES, 0.80 and 0.53 for the IFS ENS, 0.31 and 0.21 for the operational GraphCast, -0.94 and -0.96 for the GraphCast and -1.00 and -1.00 for the FuXi.

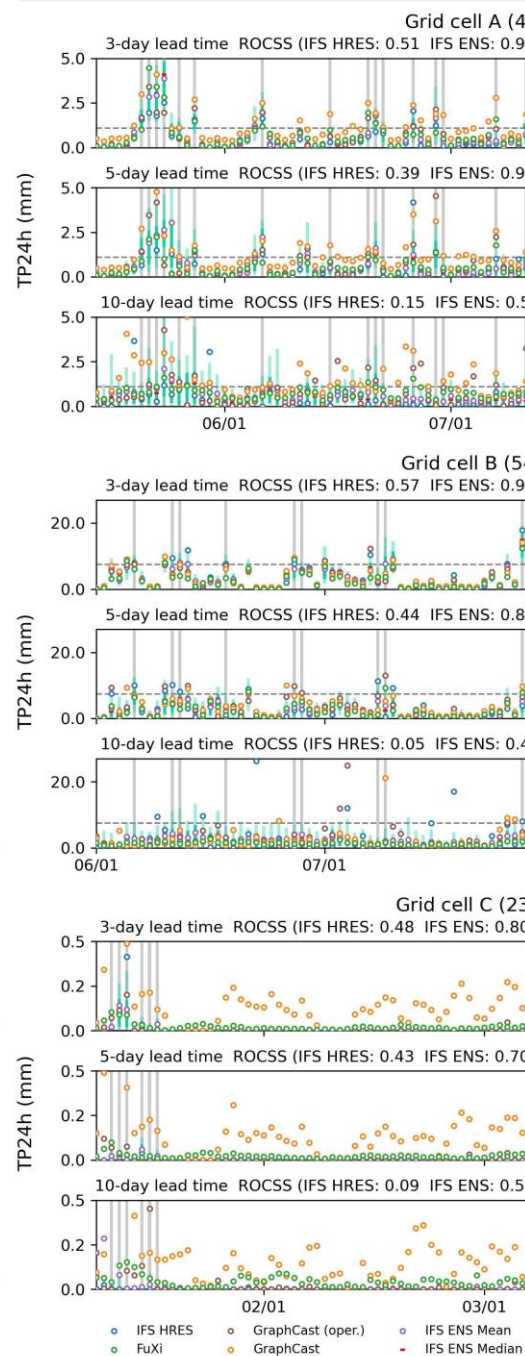
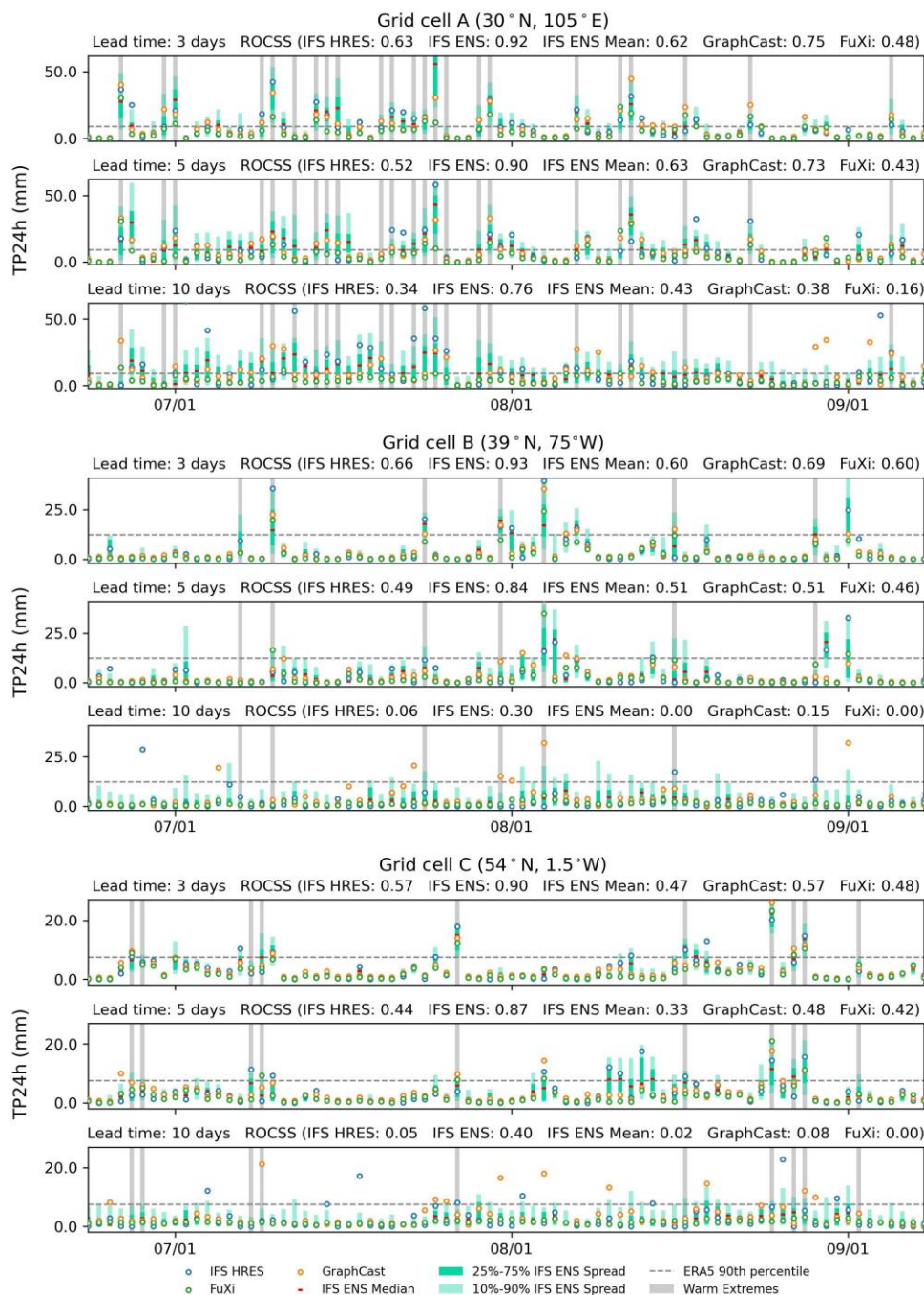


Figure 5. Time series plots of TP24h forecasts initialized at 00 UTC for the IFS HRES, IFS ENS, IFS ENS Mean, GraphCast and FuXi over three selected grid cells, i.e., A (30.44°N, 105.94°E), B (39.54°N, 75.15°W) and C (54.23.5°N, 1.520°E~~W~~).

4.3 Predictive performance of warm extremes

The differences in ROCSS for warm extremes in comparison with the IFS HRES baseline are illustrated in Figure 6. The IFS ENS tends to outperform the IFS HRES, especially in low-latitude regions. As the lead time increases, the IFS ENS tends to be more skilful than the IFS HRES. The ROCSS of the Pangu-Weather, GraphCast and FuXi is similar to that of the IFS HRES but is lower in most grids of the Pacific, Atlantic and Arctic. The GraphCast tends to outperform the IFS HRES in the Northern Atlantic near the Gulf of Mexico. The spatial patterns of the differences in ROCSS are consistent with the results of Figure 3. As the lead time increases to 10 days, the area where the Pangu-Weather, GraphCast and FuXi are more skilful than the IFS HRES decreases. On the other hand, even for lead time of 10 days, the GraphCast and FuXi continue to outperform the IFS HRES in some regions of the North Atlantic. The different performances of global weather forecasts in different regions emphasize the necessity to verify and calibrate hydroclimatic forecasts before operational application (Ben Bouallège et al., 2024; Huang et al., 2022).

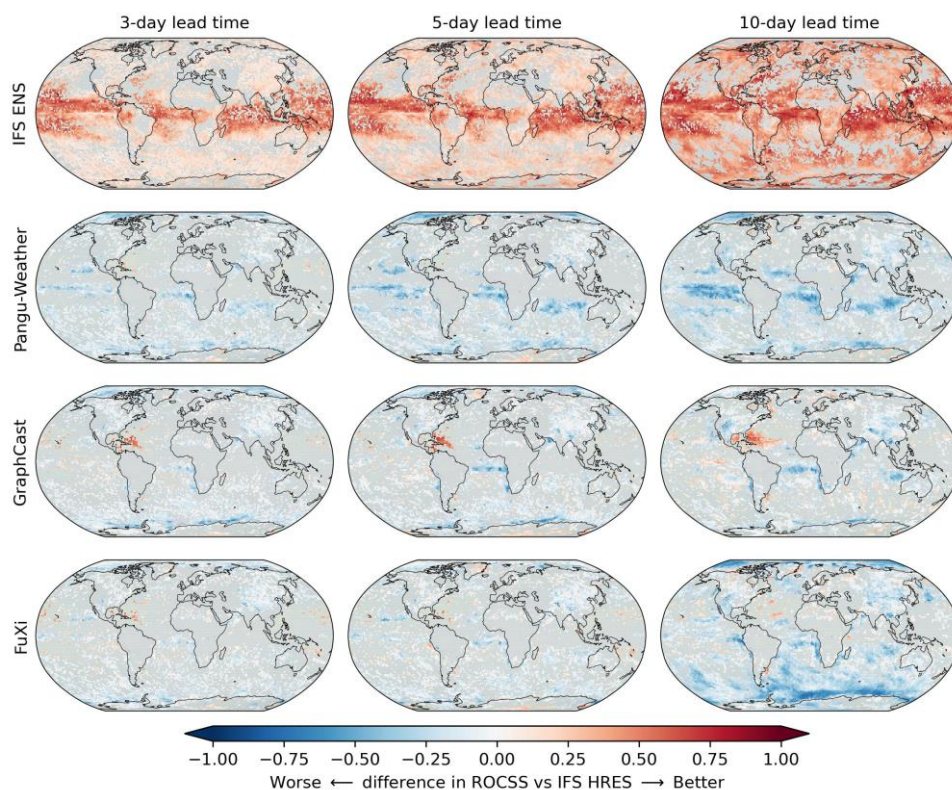


Figure 6. Differences of IFS ENS, Pangu-Weather, GraphCast and FuXi in ROCSS to the IFS HRES for warm extremes at each grid cell. The grey colour indicates grid with no statistically significant differences at the significance level of 0.1.

The time series for 24-hour maximum of 2m temperature from different forecasts initialized at 00 UTC are shown for three ~~selected~~ grid cells in Figure 7. The grid cells D, E and F are also selected respectively due to the better, close and worse performance of data-driven models in relative to the IFS HRES. Overall, the Pangu-Weather, GraphCast and FuXi exhibit similar temperature dynamics over time to those of the IFS HRES. For grid cell D, the Pangu-Weather, GraphCast and FuXi tend to outperform the IFS HRES. The Pangu-Weather tends to underestimate the temperature, leading to less true positives and more false negatives. The GraphCast and FuXi show more true positives. For grid cell E, these models show a nearly equal number of true positives and true negatives, resulting in similar ROCSS. For grid cell F, the data-driven models tend to be less accurate than the IFS HRES. The Pangu-Weather, GraphCast and FuXi tend to underestimate the temperature, leading to more false negatives and less true positives. As the lead time increases from 3 to 10 days, the ROCSS reduces from 0.48 to 0.28 for the Pangu-Weather, from 0.51 to 0.22 for the GraphCast and from 0.54 to 0.17 for the FuXi. By contrast, the IFS HRES and

IFS ENS change less. The ROCSS decreases from 0.76 to 0.56 for the IFS HRES and from 0.95 to 0.86 for the IFS ENS.

As lead time increases, these data-driven models tend to produce smoother forecasts with the skill reducing more rapidly than the IFS HRES and IFS ENS (Zhong et al., 2024; Rasp et al., 2024). For grid cells A, B and C, the IFS HRES and IFS ENS tend to outperform the Pangu-Weather, GraphCast and FuXi. At the lead time of 3 and 10 days, the ROCSS for grid cell A is respectively 0.73 and 0.42 for the IFS HRES, 0.93 and 0.86 for the IFS ENS, 0.49 and 0.24 for the Pangu-Weather, 0.51 and 0.35 for the GraphCast and 0.51 and 0.30 for the FuXi. There are more misses for the three data-driven models, indicating that they tend to underestimate the warm extremes. As the lead time increases from 3 to 10 days, the ROCSS for grid cell B reduces from 0.79 to 0.46 for the Pangu-Weather, from 0.79 to 0.46 for the GraphCast and from 0.82 to 0.51 for the FuXi. By contrast, the IFS HRES and IFS ENS change less and the ROCSS decreases from 0.81 to 0.55 for the IFS HRES and from 0.99 to 0.93 for the IFS ENS.

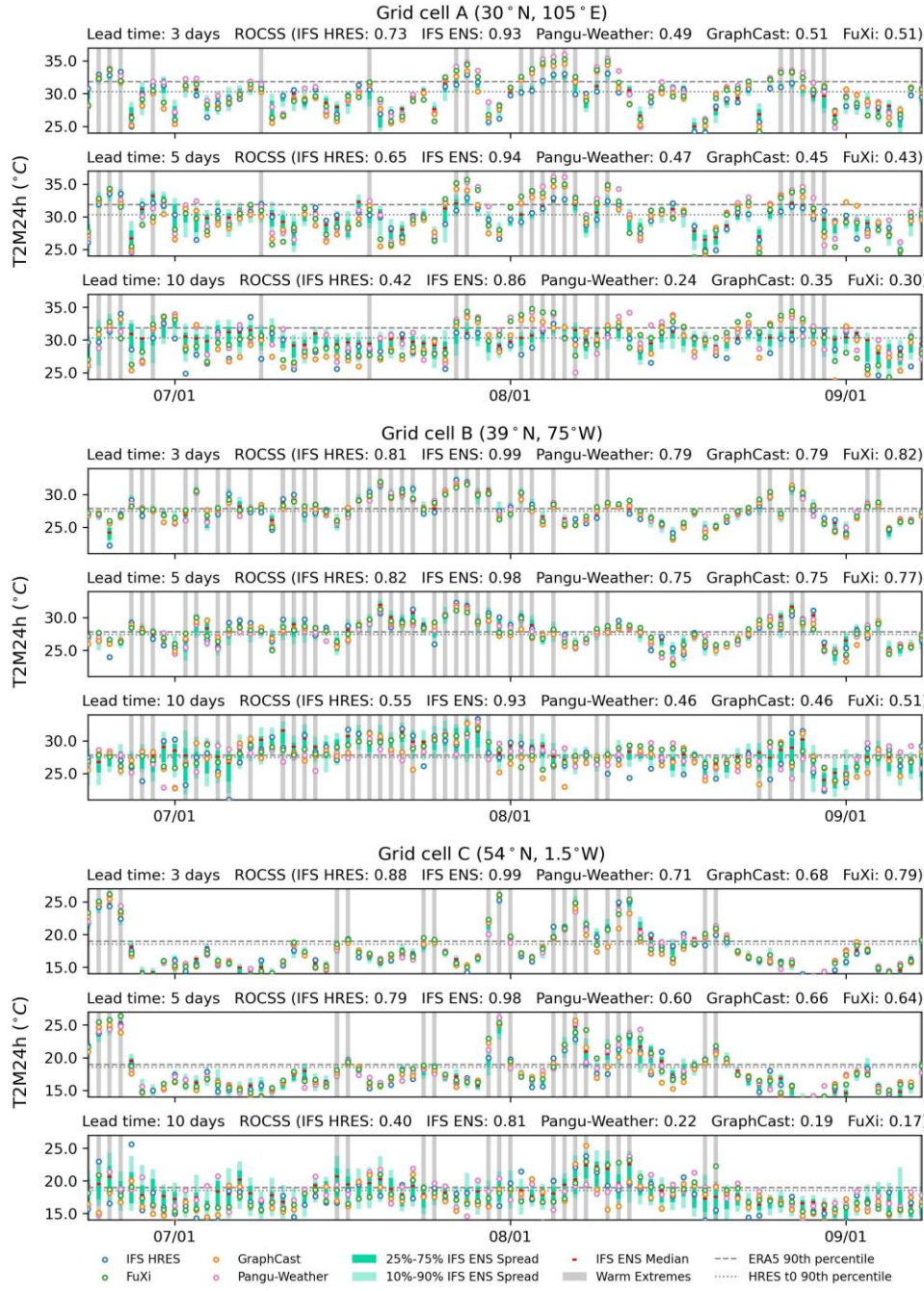


Figure 7. Time series plots of T2M24h forecasts initialized at 00 UTC for the IFS HRES, IFS ENS, Pangu-Weather, GraphCast and FuXi over three selected grid cells, i.e., **A (30° N, 105° E), B (39° N, 75° W) and C (54° N, 1.5° W) D (20° N, 75° W), E (39° N, 70° W) and F (15° S, 10° E).**

4.4 Sensitivity to predefined thresholds

The globally area-weighted performance under different predefined thresholds is illustrated for 5-day lead time in Figure 8. ~~Among the sixteen metrics, the~~ ROCSS is base-rate-independent and suitable simultaneously for deterministic and probabilistic forecasts of binary events ~~while other metrics need the predefined probability threshold to convert the probabilistic forecasts to deterministic forecasts. It is noted that the REV needs predefined cost-loss ratios to calculate the potential values of forecasts, while the cost-loss ratios may be different for hydroclimatic extremes with different threshold percentiles.~~ In the meantime, ~~it is noted that~~ the SEDI is ~~the most also~~ applicable to extreme events because of its base-rate independence and nondegenerate limit (North et al., 2013; Jolliffe and Stephenson, 2012; Brodie et al., 2024). ~~These base-rate-independent~~ metrics changes little as the predefined thresholds increase from the 80th to the 99th percentile. Specifically, as to forecast wet extremes at 5-day lead time, the scores of GraphCast decrease from 0.74 to 0.56 for SEDI and from 0.43 to 0.23 for ROCSS as the thresholds increase from the 80th to the 99th percentile. By contrast, the scores of GraphCast increase from 0.81 to 0.98 for 1-BS, from 0.87 to 0.95 for ORSS and from 0.51 to 0.52 for SEDS. These metrics are not suitable for hydroclimatic extremes because it contradicts that rarer events are often more difficult to predict (Ferro and Stephenson, 2011).

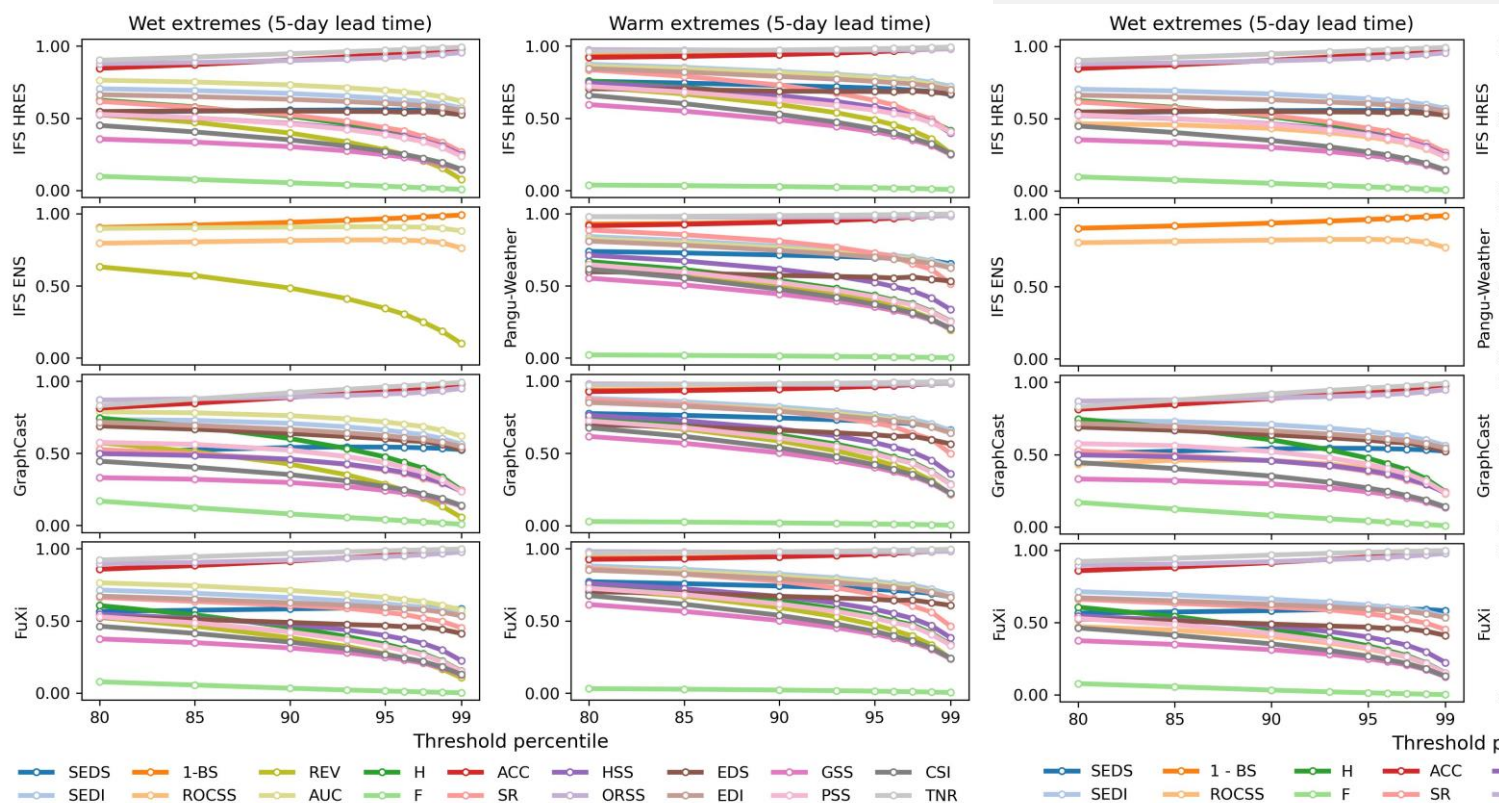


Figure 8. Globally area-weighted performance in forecasting wet extremes and warm extremes with different threshold percentiles at 5-day lead time. The REV is calculated with a fixed cost-loss-ratio of 0.2 only for purposes of illustration.

The globally area-weighted ROCSS under different predefined thresholds is shown in Figure 9. Overall, the ROCSS decreases for all eight sets of forecasts as the predefined thresholds increase from the 80th to the 99th percentile. The IFS ENS tends to perform better in forecasting wet extremes and warm extremes. Among the available data-driven models, the GraphCast (operational) tends to be more skilful for wet extremes; for warm extremes, the FuXi tends to be more skilful model at lead times less than 5 days and the GraphCast tends to be better at lead time more than 5 days. Specifically, as to forecast wet extremes at 5-day lead time, the ROCSS decreases from 0.46 to 0.24 for IFS HRES, from 0.80 to 0.77 for IFS ENS and from 0.53 to 0.26 for GraphCast (operational). As to forecast warm extremes at 5-day lead time, the ROCSS decreases from 0.69 to 0.41 for IFS HRES, from 0.93 to 0.83 for IFS ENS and from 0.70 to 0.29 for GraphCast. When the lead time is longer than 3 days, the GraphCast, GraphCast (operational) and FuXi tend to be more skilful than the Pangu-Weather and Pangu-Weather (operational) in predicting warm extremes (Olivetti and Messori, 2024b).

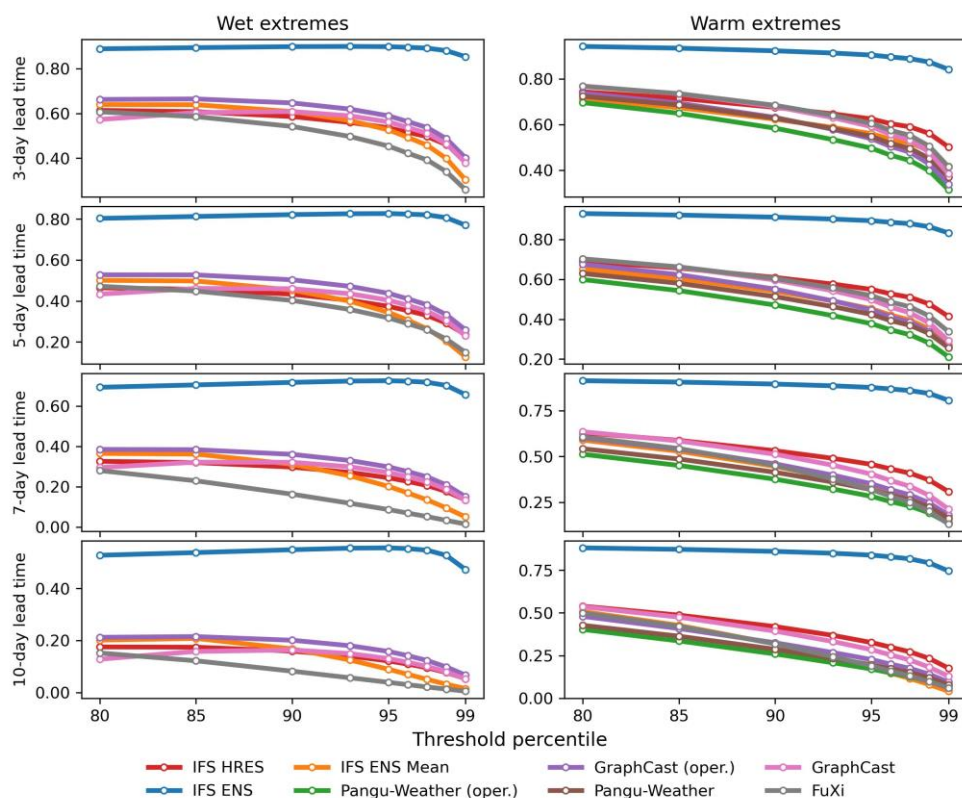


Figure 9. Globally area-weighted ROCSS for wet extremes and warm extremes with different threshold percentiles.

5 Discussion

5.1 Binary forecasts and implications on forecaster's dilemma

Binary hydroclimatic forecasts [can](#) provide useful information for disaster prevention and risk mitigation (Ben Bouallège et al., 2024; Merz et al., 2020). [The evaluation of deterministic forecasts and ensemble forecasts is usually based on Verification average metrics of deterministic and ensemble forecasts, such as the RMSE and CRPSS, that in general focus on the overall overall predictive performance, across a range of events such as the RMSE and CRPSS \(Huang and Zhao, 2022; Rasp et al., 2024\). They tend to reward Mmodels that have minimized average errors and unrealistically smooth forecasts can also be rewarded by these metrics, leading to their limited guidance to forecast hydroclimatic extremes \(Ferro and Stephenson, 2011; Rasp et al., 2020\). By contrast, verification metrics for of binary forecasts provide valuable additional](#)

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information by ~~can emphasize~~ the ability to discriminate certain hydroclimatic extremes ~~that that~~ do not directly relate to contribute little to average errors metrics (Larraondo et al., 2020). Binary forecasts are thus more suitable than continuous forecasts in these cases. In this paper, the results show that for warm extremes, the Pangu-Weather, GraphCast and FuXi tend to be more skilful than the IFS HRES within 3-day lead time but become less skilful as lead time increases. The verification of binary hydroclimatic forecasts seems to be more stringent for data-driven models since ~~the valid period of observed lead time in which~~ when there exists outperformance of data-driven models ~~are more skilful~~ is tends to be shorter than that of the previous studies for under continuous forecasts (Lam et al., 2023; Bi et al., 2023; Chen et al., 2023). In the supplement, the results across global grid cells in terms of the HSS and SEDI also support this ~~outcome result, which indicates the unique insights of binary forecasts for hydroclimatic extremes.~~ (Rasp et al., 2024) For operational applications such as disaster warning, the focus is usually on the occurrence versus non-occurrence of certain hydroclimatic extremes instead of their precise magnitude (Jolliffe and Stephenson, 2012; Larraondo et al., 2020). In the meantime, binary forecasts emphasize the ability to capture hydroclimatic extremes that contribute little to average verification metrics, ensuring that models are not rewarded for merely minimizing average errors and unrealistically smooth forecasts (Ferro and Stephenson, 2011; Rasp et al., 2020). Binary forecasts are thus more suitable than continuous forecasts in these cases. In this paper, the results show that as lead time increases, forecasts generated by data-driven models tend to be smoother and become less skilful more rapidly than the IFS HRES (Zhong et al., 2024; Rasp et al., 2024). In the meantime, ensemble forecasts can provide a range of potential hydroclimatic states and are important to quantifying the probability of hydroclimatic extremes (Price et al., 2025; Zhao et al., 2022; Pasche et al., 2025).

The hydroclimatic climate system is high-dimensional and complex so that there won't be a single verification metric to ~~determine~~ showcase all essential characteristics of a good forecast (Rasp et al., 2024; Jolliffe and Stephenson, 2012). While the verification metrics for of binary forecasts ~~can~~ emphasize the discrimination, they are unable to reflect other attributes to quantify the forecast quality, such as reliability, resolution, and uncertainty ~~and etc.~~ (SANSINE et al., 2025). ~~In the meantime, as~~ ~~shown in Fig. 5, when~~ though the GraphCast is more capable of capturing the wet extremes, it tends to produce more false positives. This result implies ~~leading to~~ the “forecaster’s dilemma”, i.e., conditioning on outcomes is incompatible with the theoretical assumptions of established forecast evaluation methods (Lerch et al., 2017). From this perspective, ~~it has been shown that using a combination of different multiple types of v~~ verification metrics and diagnostic plots ~~is in demand effective~~ (Larraondo et al., 2020; Huang and Zhao, 2022). As shown in Fig. S1 and Fig. S4 in the supplement, the values of BS for the FuXi are better than that for the HRES at the lead time of 10 days, which is different to the results for ROCSS in Fig. 4. Considering that the BS tends to reflect the average performance and is influenced by the unbalanced number of events and non-events, better values of a single metric do not mean a more useful forecast (Rasp et al., 2024). ~~Therefore~~ Overall, the

process of forecast verification needs to be guided by the demand of the operational applications and account for the trade-offs between accuracy and forecast activity (Ben Bouallège and the AIFS team, 2024; Rasp et al., 2024).

5.2 High-resolution forecasts and Use of ground truth data

High-resolution forecasts are essential for accurately capturing multi-scale processes of hydroclimatic extremes (Liu et al., 2024a; Charlton-Perez et al., 2024; Xu et al., 2025). Given the complexity of regional geographic and climatic conditions, It is noted that hydroclimatic forecasts of coarse spatial resolution tend to miss the required small-scale variability, such as the intensity and structure of typhoon (Ben Bouallège et al., 2024; Selz and Craig, 2023). In the meantime, Also, hydroclimatic forecasts of coarse temporal resolution they may might miss extreme values and the underlying evolution processes due to the mismatch between forecast time step and event time (Pasche et al., 2025). Therefore, there exists a demand to enhance the spatial and temporal resolution of data-driven models (Xu et al., 2024b; Zhong et al., 2024). It is noted that diffusion models have recently been shown to be effective for km-scale atmospheric downscaling (Mardani et al., 2025). In addition, hybrid models that utilize global forecasts from data-driven models to drive high-resolution regional models, such as the weather research and forecasting (WRF) model, can improve the forecast accuracy and resolution for extreme precipitation and tropical cyclones (Liu et al., 2024b; Xu et al., 2024b, 2025). Given that the metrics listed in Table 3 are suitable to different spatial and temporal scales, the WeatherBench 2 is capable of evaluating for high-resolution forecast data.

High-Part of forecast skill of data-driven models in-on forecasting wet and warm extremes can stem from the unfair setting of ground truth data (Rasp et al., 2024; Lam et al., 2023). As is pointed out infor the WeatherBench 2, it has beenis worthwhile to noted that the verification of precipitation using ERA5 reanalysis data as ground truth data is a compromised setting and should be considered as a placeholder for more accurate precipitation data (Rasp et al., 2024). While this comparison is not fair to the IFS models, the results indicate that using data-driven models to forecast global medium-range precipitation is promising. In addition, the verification is limited to the wet and warm extremes occurring in 2020 due to current data availability. The short verification period can only provide limited information about the model performance and sensitive results to the climate variability (Olivetti and Messori, 2024b). With the availability of As more data on hydroclimatic forecasts and baseline ground-truth observations, -and more accurate baseline data are becoming available, binary forecasts of hydroclimatic extremes deserve more in-depth verification, the capability to produce binary forecasts of hydroclimatic extremes warrants further verification. In the meantime, tThe different roles that the operational IFS analysis and ERA5 reanalysis data play in the initial conditions to generate forecasts also deserve further verification (Ben Bouallège et al., 2024; Liu et al., 2024a; Xu et al., 2024b).

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6 Conclusions

This paper has presented an extension of the WeatherBench 2 to binary hydroclimatic forecasts by utilizing ~~sixteen~~ ~~seventeen~~ verification metrics. ~~A case study is devised for binary forecasts generated by 3 data-driven models and 2 physical models.~~ Specifically, the TP24h and T2M24h are calculated from the available forecasts and ground truth in the WeatherBench 2; and the 90th percentiles of the ground truth data in 2020 are set as the predefined thresholds above which the wet and warm extremes are respectively detected. ~~Through A case study is devised for of binary forecasts generated by 3 data-driven models and 2 physical models, the results show that for wet extremes, the GraphCast and its operational version tend to outperform the IFS HRES when the total precipitation of ERA5 reanalysis data is used as the ground truth. Their globally area-weighted ROCSS is 0.46, 0.50 and 0.43 at 5-day lead time, respectively. For warm extremes, the GraphCast and FuXi tend to be more skilful than the IFS HRES within 3-day lead time while they can be less skilful as the lead time increases. At the lead times of 3 and 10 days, the ROCSS is 0.68 and 0.42 for the IFS HRES, 0.92 and 0.86 for IFS ENS, 0.63 and 0.29 for Pangu-Weather, 0.68 and 0.39 for GraphCast and 0.68 and 0.32 for FuXi. When the predefined thresholds of wet extremes increase from the 80th to 99th percentile, the ROCSS decreases from 0.46 to 0.24 for IFS HRES, from 0.80 to 0.77 for IFS ENS and from 0.53 to 0.26 for GraphCast (operational) at 5-day lead time. The extension of the WeatherBench 2 to binary forecasts facilitates more comprehensive comparisons of hydroclimatic forecasts and provides useful information for forecast applications.~~

Code and data availability

~~The forecasts and ground truth data of the WeatherBench 2 are available from the Google Cloud bucket (<https://console.cloud.google.com/storage/browser/weatherbench2>) (Rasp et al., 2024). The ERA5 data are available from the Copernicus Climate Data Store (<https://eds.climate.copernicus.eu>) at <https://doi.org/10.24381/eds.bd0915e6> (Anon, 2023a) and <https://doi.org/10.24381/eds.adbb2d47> (Anon, 2023b). A subset of the ERA5 data is also available through the WeatherBench 2 (Rasp et al., 2024). The IFS HRES, its initial conditions and the IFS ENS are available from the ECMWF's MARS archive (<https://confluence.ecmwf.int/display/UDOC/MARS+user+documentation>). The IFS ENS Mean is available through the WeatherBench 2 (Rasp et al., 2024). The training code, pre-trained parameters and access details of the data-driven models are provided in the respective papers of these models (Chen et al., 2023; Bi et al., 2023; Lam et al., 2023). The raw data, i.e., forecasts and ground truth data, used in this paper are downloaded from the WeatherBench 2 and are archived on the Zenodo under <https://doi.org/10.5281/zenodo.15066828> (Li and Zhao, 2025a) and under <https://doi.org/10.5281/zenodo.15066898> (Li and Zhao, 2025b).~~

~~The code and scripts performing all the analysis and plots are archived on the Zenodo under~~

480 <https://doi.org/10.5281/zenodo.15067282> (Li and Zhao, 2025c). All the analysis results are archived on the Zenodo under
<https://doi.org/10.5281/zenodo.15067178> (Li and Zhao, 2025d).
To guarantee future compatibility with the WeatherBench 2, the code and scripts have been made a push request to its
successor, i.e., WeatherBench-X.
The data and code to perform the complete analysis of this paper are respectively archived on the Zenodo at
485 <https://doi.org/10.5281/zenodo.14691031> and <https://doi.org/10.5281/zenodo.14691007> (Li and Zhao, 2025a, 2025b).

Author contributions

TZ: Writing – original draft, Visualization, Software, Methodology, Conceptualization. QL: Validation, Resources, Data
curation. TT: Investigation, Formal analysis. XC: Methodology, Conceptualization.

490 **Competing interest**

The authors declare that they have no known competing financial interests or personal relationships that could have
appeared to influence the work reported in this paper.

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