



Temporal and spatial scales of Near-Intertial Oscillations inferred from surface drifters.

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Abstract.

The study investigates near-inertial oscillations (NIOs) in ocean surface currents, which are important for understanding air-sea interactions and improving satellite measurements of ocean currents, such as those planned by the ODYSEA mission. Traditional methods for measuring near-surface currents, like drifters and HF radars, have limited spatial coverage, while satellite altimetry only captures geostrophic currents, missing high-frequency components like NIOs.

- Using a decade-long hourly drifter dataset (2010-2021) and outputs from the high-resolution coupled Ocean/Atmosphere model LLC2160, the study estimates the global spatial decorrelation length scales of NIOs. Drifter pairs were analyzed by latitude and distance, isolating the inertial frequency band. The LLC2160 model, which includes tidal and wind-driven forcing, was validated against the drifter data, showing good agreement in spectral characteristics and correlation scales.
- 20 Results show that NIO signals have larger decorrelation scales (~105-110 km) compared to low-frequency currents (~70-75 km), with variations by latitude (down to 50 km at high latitudes and near 20°N). These findings suggest that NIO energy is spread over larger spatial scales, linked to atmospheric forcing patterns. The LLC2160 simulation reliably reproduces these spatial characteristics and thus serves as a realistic tool for supporting ODYSEA mission planning.
 - The study shows that understanding NIO spatial coherency is critical for interpreting satellite Doppler measurements and mitigating aliasing effects, thereby improving ocean current observation capabilities.

1 Introduction

Surface and near-surface currents play an important role in many processes and applications through air-sea interactions (Renault et al., 2016, 2024, Nuijens et al., 2024), transport of drifting material (Onink et al., 2019, van Sebille et al., 2019), or their impact on sea state (Ardhuin et al., 2017). Today measurements of near surface currents are performed with High-Frequency coastal radars (Roarty et al., 2019) or in situ drifters (Lumpkin et al., 2017, Elipot et al., 2016), both with very limited or sparse spatial coverage. These measurements are complemented by indirect estimates of surface currents based on

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sea level anomalies using satellite altimetry, combined with an estimate of the mean dynamic topography that also uses a climatology of drifter currents (Mulet et al., 2021). Because these altimeter-derived "geostrophic currents" miss a significant part of the "total surface current", alternative methods are being developed to map surface currents more completely. One approach might be to increase the number of drifters in the ocean. Another promising solution is the perspective of using Doppler radars on satellites (Rodriguez 2018, Ardhuin et al., 2019, Wineteer et al., 2020). Although satellites can provide unprecedented spatial coverage, they can only provide snapshots of the surface current field that contain both low-frequency currents, mostly in geostrophic balance away from the Equator and thus measurable from satellite altimeters, and highfrequency currents with a 1 ominant contribution from the near-inertial frequency band (Yu et al., 2019). Although a longterm average of a sequence of Doppler data can provide useful information, the most novel aspect of this 2ew data is their potential to document the fast-varying and unbalanced components. Early analysis by Ubelmann et al., 2021 suggested that measurements across a 300-km wide swath with a 3-day revisit at mid-latitudes could provide enough constraints to recover more than half of the energy of 4ear Inertial Oscillations. 3hat analysis was based on modeled currents underestimated NIO velocities and critically relied on the spatial coherence of the NIO velocity field. To better understand the representativeness of in situ or satellite data it is therefore critical to understand the spatial structures of NIOs from 5 tual observations. 6 Ithough NIOs are generated by winds and are thus expected to contain some large scale pattern (O(1000km)), associated with the large atmospheric Rossby radius. $\overline{7}$ hey also respond to the full spectrum of wind patterns and are also known to interact with the ocean mesoscale field potentially resulting in local change of phase and amplitudes.

The ODYSEA (Ocean Dynamics and Sea Exchanges with the Atmosphere) concept mission is a recent satellite mission proposing to measure surface currents and wind stress using a Doppler Scatterometer, over a swath and with a temporal resolution of about 12h at mid-latitudes (Rodríguez et al., 2019; Wineteer, et al., 2020; Torres et al., 2023). The 12h revisit period is not short enough to resolve the seriod of inertia at mid-latitudes, and the NIO signal pay alias the ODYSEA measurements (Wang et al., 2023). Topper characterization of these NIO signals is crucial for a satellite mission like ODYSEA. The aim of this study is therefore to propose a method for estimating a characteristic decorrelation length of the NIO signal in the ocean from the available drifter database. A second objective, in support of future studies for ODYSEA, is to verify that these spatial correlation scales are consistent with simulations from Ocean/Atmosphere coupled models.

To answer these questions, section 2 of this article presents the data from model and in situ and the implemented to estimate the decorrelation length of the NIO signal. The high frequency ocean surface currents are characterized through the analysis of ocean drifting buoy velocities.

Section 3 presents the results obtained from the model and in situ observations focusing on the NIO through their spatial coherency and 12 relation; and finally, section 4 details the conclusions of this study.

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I would nuance this statement based on latitude, basin, and the actual frequency range of what you define as as "high-frequency". Look at Figure 2 of Yu et al.
2019 or Figure 2 of Arbic et al. 2022.
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Number: 3 Author: Subject: Comment on Text Date: 8/19/25, 12:05:11□PM
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Number: 4 Author: Subject: Comment on Text Date: 8/19/25, 12:04:42 PM
I think you want to define the acronym NIO here?
Number: 5 Author: Subject: Comment on Text Date: 8/19/25, 12:07:36□PM
What are "actual observations" with respect to "in situ or satellite data". Are they different? I am not following.
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NIOs at the surface yes, but I think spontaneous inertial oscillations could be also generated by current instabilities.
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Not sure what you are trying to say here. That they are not uniform on the scale of the atmospheric forcing?
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In which way? Could you explain how getting the decorrelation time scale of NIOs will help with the aliasing issue? Does it help?
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correlation of NIOs? What is that?





2 Data and Method

2.1 Hourly Surface Velocity Program Drifters

(http://www.aoml.noaa.gov/phod/dac/dacdata.php) and described in Elipot et al., 2016 is used. Both velocity at 15m (when drifters are drogued) and at the surface (when the drogue is lost) are provided. The analysis of the high frequency content of this hourly dataset, and it is placed and zonally averaged anticyclonic and cyclonic rotary spectra of drifter velocities, show that the NIO frequencies can be retrieved. This database has been used extensively to study hight frequency. ocean signals, and has the estimation of the wind power on NIO (Liu et al., 2019), he near inertial current (Chaigneau et al., 2008) or dulerian versus Langrangian currents (Caspar-Cohen et al., 2022; Zhang et al., 2024).

For this study, the amount of data available in the hourly database is low before 2004. So, for computational efficiency and cost, we have limited the time frame of our analysis to a period ranging from 2010 to 2021, which contains lost of the data. Both surface (undrogued) and 15m (drogued) drifters are used, but they are analysed separately because and lost impacting low frequency of the signal (Etienne et al., 2024). We have thus decided to keep the observed surface velocity uncorrected for this study.

Hourly drifter velocity dataset distributed by the Global Drifter Program (GDP) Data Assembly Center (DAC)

2.2 LLC2160-C1440

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The outputs of surface current from a numerical model are considered for the statistical comparison of NIO characteristics with what is obtained from the drifters. We chose a simulation from an Ocean/Atmosphere coupled models which should be prone to represent realistic levels of energy for the wind driven current and the NIOs in particular.

The LLC2160-C1440 (Torres et al., 2022 and hereafter LLC2160) simulation is based on the ocean component of the Coupled Ocean-Atmosphere Simulation (COAS) using the Goddard Earth Observing System (GEOS) atmospheric and land model coupled to an ocean configuration of the Massachusetts Institute of Technology general circulation model (MITgcm). The description of the GEOS model can be found in Molod et al., (2015) and Strobach et al. (2020). MITgcm is based on the PatitudeLongitude-polar Cap 2160 (LLC2160) configuration described in Arbic et al. (2018). Horizontal grid is of 2–5 km resolution over the ocean (global horizontal grid spacing of 1/24°) and 7 km for the atmosphere. The vertical grid has 90 levels with thickness ranging from 1m near the surface to 300m below 5000m. This ocean model configuration includes tidal forcing and two-way coupling between winds and ocean currents on wind stresses (Torres et al., 2022) to improve wind and current variability at short time scales. We use here the surface current outputs at hourly frequency of a year of a year of 10 climatological velocity fields. Since both the Adag sphere and the Oda have a free evolution, we do not expect any particular match between the phases of the drifter and simulation 14 and, so the comparison will be purely

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2 atistical. 1 or this reason, it was reasonable to consider the run as a statistical realization repeated every year over the longer drifter time extension.

LLC2160 hourly surface data are interpolated onto drifter time and positions whatever the status of the drogue surface or 15m motion) to provide "Lagrangian-like" LLC2160 trajectories. Hence, in the following, surface currents from LLC2160 are compared to surface and 15m currents derived from drifters, with respect to the spatial and temporal sampling of these observations.

In a recent study, Zhang et al., (2024) demonstrated that in the near-inertial spectral band the Lagangian and Eulerian kinetic energy (KE) levels are comparable, while Lagrangian KE can underestimate Eulerian KE at tidal and low frequencies. Hence, using" Lagrangian-like" estimation of the velocity may not affect the LLC2160 comparison to Lagrangian drifter displacement.

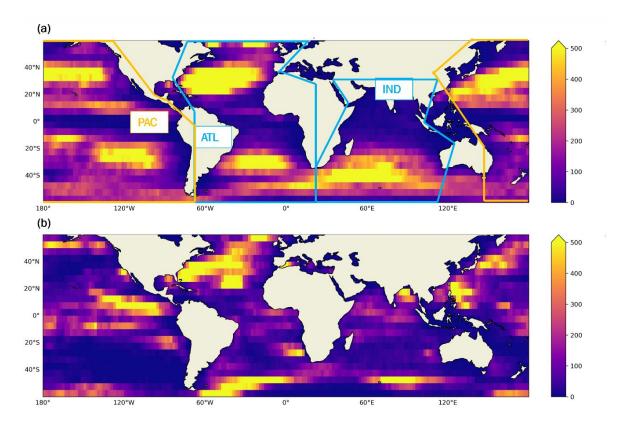
105 2.3 4 election of pair of drifters and preprocessing

To select pairs of drifters in the near-inertial frequencies, we select drifters in 5° latitude bins. This makes it possible to obtain trajectories whose time series contain similar sertia frequencies, which depends on latitude. We then select pairs of buoys that have more than 3 days overlap and isolate the inertial signal, whose periods range from more than days near the equator to 12 hours at high latitudes. Jajays is a compromise between the number of selected pairs (vanishing when overlapping period increases) and the ability to the Fourier transform tools to separate the entire inertia the quator. Hence a low bound of 3 days for the selected time series allows to study NIO excluding only the 120° around the equator. When pairs are identified, distance in kilometer that we the drifters in a pair is computed. The total number of pairs of drifters selected in 5° latitude bins are shown in Fig.1. Surface (undrogued) and 15m data (drogued) drifters will sample different dynamics, as the spatial distribution of the two datasets is different and the surface motion more sensitive to the wind forcing (and more generally to the atmospheric conditions). 14 sically, undrogued drifters are mainly trapped into gyres, while 15m drifters follow the main currents. Hence, some basins are sparsely sampled by the drifters, as the Indian ocean at the surface or south Pacific at 15m (IND and PAC respectively on Fig.1).

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	d say it is more your method than the Fourier transform in itself.
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, 0	ne inertial period is 2.8 days. I think you should rewrite this part as it is a bit unclear.
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	d does not take into account the release patterns of the drifters. However, it a true that the density of undrogued drifter observations is higher es of the general circulation.
convergence zone	so of the general enoughtern.







120 December 2021. Data are binned into 2° longitude x 5° latitude. Regional areas are delimited in blue (ATL for Atlantic Ocean, IND for Indian Ocean) and in orange (PAC for Pacific Ocean).

2.4 Bandwidth filtering

frequencies below 0.9f from the total signal.

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The inertial period T_i = 2π/f depends on the inverse of the Coriolis frequency f. T_i tends toward infinity near the equator and toward 12h at high latitudes to the higher is the latitude, the higher is the inertial frequency. We chose to apply a Lanczos filter (Duchon 1979) on the drifter trajectories to extract the inertial signal. A bandwidth selection of {0.9f – 1.1f} is applied for each 5° latitude binned drifters of the pairs, with f computed for the mean latitude of the 5° bin. An example of such filter is shown in Fig.2 for two surface drifters of a pair, numbered 1820 in the Pacific Ocean between 45°N and 50°N. Trajectories of the drifters are close together and show a strong NIO event from explored of January to mid-February 2015.

Chertia loops are clearly seen on Fig.2 (a) and NIO signals are have during this event (Fig.2 (c)).

In the following, we study both the NIO signal and what we called the "low" frequencies, computed by removing

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towards
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obvious and not needed. You can just write f = 2 \Omega \sin latitude.
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tends to 12h towards the poles?
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inertial not inertia
Number: 7 Author: Subject: Comment on Text Date: 8/19/25, 1:01:36□PM do you mean "in phase" across the two drifters?
do you mean in phase across the two differs:





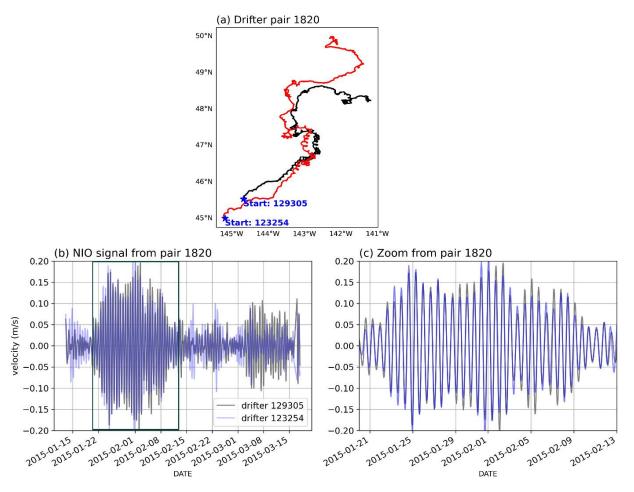


Figure 2: Example of pair numbered 1820 selected between 45° N and 50° N in the Pacific. (a) shows trajectories of the two drifters 129305 and 123254. Times series of the NIO extracted from the zonal velocity on drifter 129305 (black) and 123254 (blue) are shown in (b), with a zoom in the black rectangle shown in (c).

2.5 Rotary spectra and correlation between pairs

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The globally averaged anticyclonic and cyclonic rotary spectra are computed based on the Elipot et al., (2016) method: trajectories are split into segments of 2160 hours (2ear 90 days) with 50% overlapping. Then the discrete Fourier transform of the complex velocity fields (u+iv, where u and v are zonal and meridional velocity, respectively) is computed. In the following, the rotary spectra are calculated using complete buoy trajectories (and not the selected pairs).

Anticyclonic frequencies (negative in the northern hemisphere, positive in the southern hemisphere) and cyclonic frequencies (with the opposite sign of anticyclonic) are averaged to obtain a global value. The same method is applied by averaging in 5° latitude bins to observe latitudinal dependencies of the signal variability.

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exactly 90 days?		





Different diagnostics and statistics based on drifter pairs have been estimated and compared to evaluate their relevance in the estimation of the NIO spatial coherency. We retain here the correlation between drifters of a pair binned by 5km distance. Hence, for each drifter (x,y) of a pair p in a distance bin d, anomaly is computed by removing their respective mean value x_d^p and used to compute the covariance $COV_d(x,y)$, variance $VAR_d(x)$ to correlation $COR_d(x,y)$:

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$$x_d^p = \frac{1}{Np} \sum_{i=1}^{Np} x_i^p \tag{1}$$

$$COV_d(x, y) = \frac{1}{N} \sum_{p=1}^{P} \sum_{i=1}^{Np} (x_i^p - x_d^p) (y_i^p - y_d^p)$$
 (2)

$$VAR_d(x) = \frac{1}{N} \sum_{p=1}^{p} \sum_{i=1}^{Np} (x_i^p - x_d^p)^2$$
(3)

$$COR_d(x, y) = \frac{COV_d(x, y)}{VAR_d(x)^{1/2}VAR_d(y)^{1/2}}$$
(4)

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Where N_p the length of the selected time series of pair p in the distance bin d and $N = \sum_{p=1}^{P} N_p$ the total number of data point in the distance bin (sum over the P existing pairs). Hence, we can compute the correlation between drifters of the pairs as a function of their distance, but also as a function of latitude. In this case, pairs are selected in $\pm 2^{\circ}$ around the specified latitude (corresponding to a 5° latitude band) and statistics computed very a 1°.

160 3 Results

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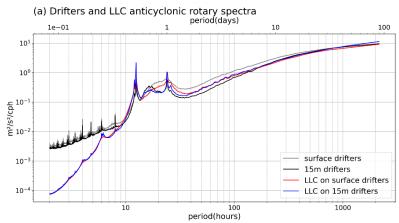
3.1 Assessment of LLC

A first question to be desolved is to which extent the LLC2160 is realistically representing sertial current statistics. Global rotary spectra computed separately on 15m and surface drifters show a better match between surface drifter and LLC2160 (as mentioned in 2.2, only surface currents from LLC2160 are interpolated on surface and 15m drifter data). On surface and 15m anticyclonic rotary spectra are shown. The cyclonic rotary spectrum is not shown here as he NIO is only an anticyclonic motion. Energy level in the surface drifters (grey curves in Fig.3) is systematically higher than at 15m (black curves in Fig.3). Exception can be noticed in the large maximum of surface spectrum between approximately 16h and 30h periods, which is narrower in the 15m rotary spectra. This local flat distribution is the result of the wind-driven inertial oscillations, which is a combination of the variation in the sertia frequency with latitude and the latitudes where the wind is predominant (Elipot and Lumpkin, 2008; Chaigneau et al., 2008). The effect of wind diminishes with depth, the energy associated with this process should be lower at 15m than at the surface.

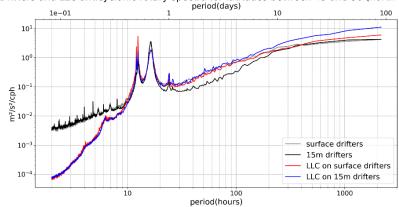
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as NIOs are assumed	d to be solely anticyclonic.	
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		II. 2022 (Figure 2 with version 1.04c of the hourly dataset) where the globally averaged spectrum of drogued ters for a narrow range of frequencies within the global inertial range. I am wondering if this has more to do
9	stribution of the drogued vs undr	
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Why is this an explan	nation for the black spectrum to b	pe higher than the gray spectrum? I am not following.



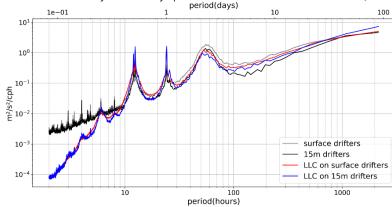




(b) Drifters and LLC anticyclonic rotary spectra for latitudes between 45 and 50 (north and south)



(c) Drifters and LLC anticyclonic rotary spectra for latitudes between 10 and 15 (north and south)



1 igure 3: Anticyclonic rotary spectra of surface drifters (grey lines), 15 m drifters (black line), LLC2160 collocated on surface drifters (red) and on 15 m drifters (blue). The global rotary spectra in (a) are the average rotary spectra over the entire domain. (b) are the average rotary spectra between $\pm 45^{\circ}$ and $\pm 50^{\circ}$. (c) are the averaged rotary spectra between ± 10 and $\pm 15^{\circ}$. Period is expressed in hour (lower axis) and day (upper axis).

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In subplots b and c it would be useful to show the inertial frequencies: f in the middle of the bin, 0.9f, 1.1f, and the inertial range of the bin.



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The global difference in the surface and 15m drifter spatial and temporal distribution (Fig.1) partly explains this exception between the drifter rotary spectra. This point is supported by the LLC2160 spectra, that also differ while the same surface velocity fields (same dynamics) are used for both curves (red and blue in Fig.3).

At short periods (below 10h), the LLC2160 dramatically underestimates the drifter rotary spectra, but this represents only a small proportion of total energy. From the 12h 2ic maximum and the unit period of one day), LLC2160 and drifter curves show the trend at similar energy level, even if LLC2160 hergy is generally weaker than the drifter one (for each distribution respectively). As the minimum period of the inertia 12 hours (at high latitudes), we have suppose the LLC2160 resolves quite well the NIO frequencies.

The main tidal energy pies 10 day and 12h) are a bit overestimated by the LLC2160. Out as stated in Zhang et al., (2024), at tidal frequency Lagrangian spectra is smoother than the Eulerian counterpart, and this result is not critical for this study.

Above 1 day period, the LLC2160 spectra lie between 111 face and 15m drifter curves 12 owing an underestimation (overestimation) of the surface (15m) drifter signal respectively. It 13 kes sense considering that the model does not include surface Stokes drift 14 at nearly disappear in the 15m drifters) but includes a stronger Ekman component than at 15m.

Finally, LLC2160 low frequency signal (i.e. periods greater than 10 days) fairly match the drifter spectra.

If we focus on narrower latitude bands, where the range of NIO frequency is well defined, as in Fig.3 (b) and (c), the inertial peak becomes sharper and LLC2160 still fits quite well the drifter rotary spectra for these frequencies. Between $\pm 45^{\circ}$ and $\pm 50^{\circ}$, the inertial period is near 18h (Fig.3 (b)) and near 58h between ± 10 and $\pm 15^{\circ}$ (Fig.3 (c)). In both cases (and for both 15m and surface distributions), the inertial maximum is clearly displayed in the LLC2160 but remain slightly weaker compared to drifters (by approximately 20%). Finally, we also see a better match between LLC2160 and the surface drifters compared to 15m drifters, especially at low frequencies.

The rotary spectra computed as a function of latitude on Fig.4 brings an interesting global view of the near inertial energy, localized along the white-dotted line. As expected by the integrated spectra of Fig.3, 15 inertial energy of the LLC2160 simulation is quite consistent with the LLC2160 simulation globally at all latitudes. However, we note a slight underestimation of inertial energy by the LLC2160 simulation a bit more pronounced around 40°N and near the equator. As can be seen on both Fig.3 and Fig.4, the sub-inertial frequency power spectrum is strongly reduced from the surface to 15m, 16 the surface drifters are submitted to the atmospheric boundary layer with the direct interaction with the wind (surface Ekman layer, waves), while this interaction is decreasing with depth (i.e. Ekman spiral, Stokes drift etc...)

For this reason, we will consider in the following only surface statistics.

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What? LLC2160 is consistent with LLC2160?			
Number: 16	Author: Subject: Comment of	•	
Would probably be e	asier to say that atmospheric-dri	ven current variability typically decreases away from the surface.	





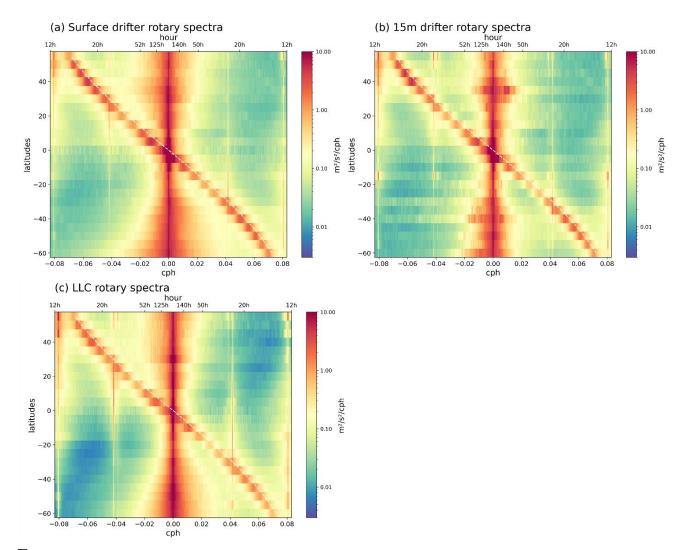


figure 4: Rotary spectra of surface drifters (a), 15 m drifters (b) and collocated LLC2160 (c) on surface drifters as a function of latitude. Rotary spectra are averaged in 5° latitude bins. Frequency is expressed in cycle per hour (cph) on the bottom axis and respective period in hour is indicated in the top axis. Inertial period is indicated by the white line.

3.2 NIO spatial correlation scale

215

Correlation between drifter pairs is evaluated following Eq. (4) in 5km bins, using surface drifter atasets. We also show the statistics of the low frequency signal (< 0.9f), extracted from the filtering process as described in 2.4.

The curves of the mean global correlations computed on the urface zonal velocity feature similar characteristics in the drifters and the simulation (Fig.5 (a)). Both low frequencies and NIO correlations match the drifter statistics. The 0.5 threshold for the correlation, that we call the surface decorrelation distance on Fig.5 (b), is reached at 105-110 km for the

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grammar: "the surface drifter dataset."			
Number: 3 Author:	Subject: Comment on Text	Date: 8/19/25, 2:22:27 PM	
Why limiting yourself to zonal velocities? Since inertial oscillations are circular in nature, do you get the same results?			



225



NIO, while it is only 70-75 km for the low frequencies. This is an interesting result suggesting that the near-inertial energy is mainly distributed over the wide spatial scales, probably related to the size of atmospheric fronts at first order. At a second order, there are known interactions between mesoscales and NIO (e.g. Young and Ben-Jelloul, 1997) but they are certainly not dominant. The NIO energy would remain dominant at the larger scales that of the mesoscales.

However, for distances longer than 300km, the correlations reach a threshold, around 0.2 for low frequencies and 0 for NIO. This might be explained because low frequencies contain multiple dynamics at many different ranges of spatial scales, 2 cluding very large ones remaining correlated far beyond 300km.

The surface decorrelation distance from the drifters has a 3patial variability (latitude dependency) that LLC2160 correctly reproduces (Fig.5 (b)). Decorrelation distance of the NIO signal at the surface can reach more than 150 km near the tropics, down to 50 km at high latitudes (near $\pm 60^{\circ}$) and at 18°N. We do not have a clear explanation for these zonal variations, probably related to the size of the atmospheric patterns responsible for the inertial energy.

For the low frequencies, the match between LLC2160 and the drifters is also very good. We would expect the decorrelation distance to be related to the Rossby radius size, 4 hich seems true from 30°N to 60°N, but not as clear elsewhere.

Over most of the global ocean, NIO and low-frequency signals have different decorrelation scales, which can be used to differentiate and separate them.

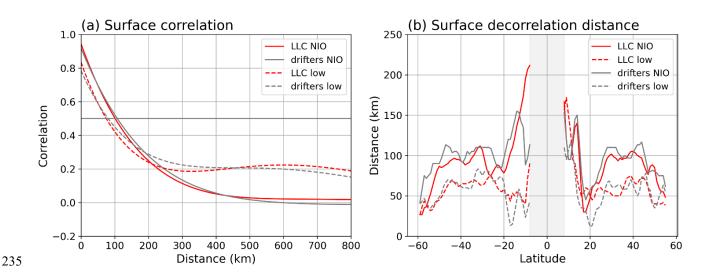


Figure 5: global surface zonal velocity correlation (a) and surface decorrelation distance (distance at the correlation threshold of 0.5) as a function of latitude (b). Drifter values are plotted in grey and collocated LLC in red. Solid and dash lines indicate values for NIO and low frequency respectively. In the $\pm 8^{\circ}$ region, delimited by the shaded area in (b), signal cannot be properly filtered.

240

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Since you are limiting yourself to latitudinal dependency, which states this.		
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Can you bring in values from published studies?		
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4 Conclusions

245

This study ① characterizes the spatial decorrelation scales of near-inertial oscillations (NIOs) in ocean surface currents using both drifter observations and the LLC2160 coupled ocean-atmosphere model. ② have set up a method for characterizing NIO correlation scales. The spatial coherency of the NIO as seen at the surface has been investigated using a ③ o years hourly SVP dataset. By selecting pairs of drifters binned by latitude and distance, we have estimated the decorrelation length scale ④ f the NIO, together with the low frequency signal (<0.9f).

The results show that NIOs exhibit larger spatial coherence (~100 km on average) compared to low-frequency currents (Fig.

5 (a)), with significant latitude-dependent variability. The LLC2160 model effectively reproduces the observed NIO spatial patterns and energy spectra, supporting its use in anticipating challenges such as signal aliasing in future satellite missions like ODYSEA, which aims to measure surface currents and wind stress using a Doppler Scatterometer with a 12-hour revisit at mid-latitudes. Thus, LLC2160 constitutes a realistic reference been for the ongoing ODYSEA studies. For especifically, as suggested in a companion study, the large scales of the NIO patterns can help mitigating aliasing of the ODYSEA signal (S. Jousset personal communication).

Data availability:

255

Hourly drifter observations are produced by Atlantic Oceanographic and Meteorological Laboratory (AOM) of the National Oceanographic and Atmospheric Administration (NOAA). Drifter data can be **Zownloaded** from tp://ftp.aoml.noaa.gov/pub/phod/buoydata/

260 The LLC2160 simulation over year 2020 can be found at https://portal.nccs.nasa.gov/datashare/G5NR/DYAMONDv2/GEOS_6km_Atmosphere-MITgcm_4km_Ocean-Coupled/GEOSgcm_output/geosgcm_surf/

Author contribution:

HE carried out the study and diagnostics. CU, FA and GD supervised the study. HE prepared the article with contributions from CU and FA.

Competing interests:

The authors declare that they have no conflict of interest.

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Disclaimer

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