



Brown carbon emissions from laboratory combustion of Eurasian arctic-boreal and South African savanna biomass

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Abstract:

Warming climate is predicted to increase forest fires which can be a major source of black and brown carbon (BC and BrC) into the atmosphere. Unlike North American forest fires, very limited studies have characterized North Eurasian biomass burning (BB) emissions. In this work, we defined the emission factors of carbonaceous aerosols and characterized light absorption of BrC emitted from boreal and peat burning through offline filter extraction method. The results were compared to African savanna emissions. Effects of atmospheric dilution and oxidative aging on BrC absorptivity were investigated for selected BB emissions sampled into an environmental chamber. Organic carbon (OC) and elemental carbon (EC) emission factors of fresh BB emissions ranged between 1.30–89.9 g kg⁻¹ and 0.01–4.80 g kg⁻¹ respectively. Methanol soluble OC (MSOC) represented more than 92% of fresh BB emissions but intrinsic chemical differences among samples resulted in MAC_{MSOC} values ranging from 0.46–1.48 m² g⁻¹ at 365 nm. Fresh BB emissions formed weakly absorbing BrC with k_{550_MSOC} ranging from 0.002 and 0.011. Water soluble OC (WSOC) fractions varied among fresh BB emissions but overall exhibited higher MAC₃₆₅ than MSOC. Dilution-related



38 evaporative loss in environmental chamber resulted in less volatile OC, making them less soluble in
39 methanol. Photochemical and dark oxidative aging further increased the ELVOC fraction of the organics
40 along with oxidation state. Our estimated OC-EC emission factors and k_{MSOC} for fresh BB emissions can be
41 used for future modelling purposes. Further online measurements are needed to account for non-soluble
42 strong BrC in aged BB emissions.

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45 **Introduction:**

46 Biomass burning (BB) emission is one of the largest anthropogenic sources of black (BC) and brown carbon
47 (BrC) in the atmosphere (Bond et al., 2004; Kirchstetter et al., 2004). Multiple studies have predicted that
48 open BB emissions such as wildfires will become more prevalent in the future with the warming climate,
49 with increasing boreal forest and peatland fires threatening also the Eurasian regions (Krawchuk et al., 2009;
50 Costa et al., 2020; Feyen et al., 2020). Along with the expected overall increase in the frequency of “high-to-
51 extreme” wildfires in Europe (de Rigo et al., 2017) an expected shift in vegetation distribution from Southern
52 towards Northern Europe might add stress on Finno-Scandinavian biomass and amplify its fire hazard (Costa
53 et al., 2020). Only limited laboratory and field studies exist describing aerosol emissions from boreal forest
54 and peatland fires in the Eurasian area (Wilson et al., 2015; McCarty et al., 2021; Zhong et al., 2024 ;
55 Schneider et al., 2024a), compared to the more studied North-American boreal forest and peat fires (Aurell
56 and Gullett, 2013; Urbanski, 2013; Stockwell et al., 2014; Black et al., 2016; Andreae, 2019; Phillips et al.,
57 2022; Zhao et al., 2021). However, the fire regimes in Eurasia are known to be dominated by surface
58 burning, that likely emits carbonaceous aerosol of different characteristics in comparison to more crown-fire
59 dominated fires in North America (De Groot et al., 2013; Rogers et al., 2015).

60 Chemical and optical properties of BC have been studied extensively over the last two
61 decades and it has been established as an important climate warming agent (Jacobson, 2000, 2001; Bond et
62 al., 2013). On the other hand, the properties of combustion emitted BrC have still not been fully
63 characterized due to variability in combustion conditions (Martinsson et al., 2015; Wang et al., 2020; Saleh,
64 2020), fuel chemical composition (Saleh et al., 2014; Smith et al., 2020; Moschos et al., 2024; Navinya et al.,



2024) and secondary transformation of BrC in the atmosphere (Laskin et al., 2015; Brown et al., 2018; Hems et al., 2021). BB derived BrC consists of different light-absorbing organic precursors of BC, such as polyaromatic molecules, which are not transformed to fully ordered BC during the combustion process (Saleh et al., 2018). Since BC formation is strongly temperature-dependent, low combustion temperatures may favour BrC formation, whereas high combustion temperatures increase oligomerization/polymerization of polyaromatic carbon structures to form BC (Faccineto et al., 2011; Desgroux et al., 2013; Solum et al., 2001; Wang, 2011). Consequently, the light-absorbing primary organic aerosols (POA) emitted from BB emission are chemically diverse and distinct for different fuel types and combustion conditions. Additionally, they evolve in the atmosphere through oxidation and functionalization reactions, driven by the formation of secondary organic aerosol (SOA) and the evaporation and chemical fragmentation of organic aerosol (OA), resulting in a complex mixture of organic chromophores with variable absorptivity.

BB is a dynamic and variable process that strongly depends on fuel moisture content, fuel composition and combustion conditions. A common parameter to characterize BB is the modified combustion efficiency (MCE), which describes the share of carbon dioxide to the sum of CO and CO₂ emissions. However, studies have typically found only weak correlations between BrC light absorption and MCE (Pokhrel et al., 2016; McMeeking et al., 2014). The ratio of the emitted BC (mainly composed of elemental carbon or EC structures) to the organic carbon (OC) has been suggested as a more suitable parameter for correlating the aerosol optical properties with combustion conditions (McClure et al., 2020; Stockwell et al., 2016; Saleh, 2020). For instance, flaming dominated combustion processes generally lead to relatively high EC/OC ratios and more absorptive BrC than smouldering emissions (Saleh et al., 2014; Xie et al., 2018; McClure et al., 2020; Kumar et al., 2018).

In climate models, the light absorption strength of the material is described by the imaginary part of the refractive index (k). For pure BC, k - are constrained close to unity and exhibit very little to no wavelength dependence, especially at UV and shorter visible wavelengths (Bond and Bergstrom, 2006). However, the k for BB emitted BrC seem to vary across several orders of magnitude depending on the biomass type and combustion conditions as well as the measurement techniques (Chakrabarty et al., 2010; Bluvshstein et al., 2017; Saleh et al., 2018; Saleh et al., 2020; Navinya et al., 2024). Therefore, characterizing



the k for individual BrC compounds is a daunting and, in some cases, impossible task due to instrumental limitations. A more comprehensive approach has been adapted in the last decade where the k values for BB derived BrC has been shown to fall in a “brown-black carbon continuum” (Saleh et al. 2018). BB derived BrC generally exhibit progressively higher k values with increasing temperature, and has been termed as “dark BrC (d-BrC)” (Hoffer et al., 2017; Adler et al., 2019; Atwi et al., 2022; Chakrabarty et al., 2023) or “strongly absorptive BrC (s-BrC)” (Mclure et al., 2020; Saleh et al., 2020) when the combustion temperatures approach the BC formation regime. In contrast, the BrC derived from BB combustions in progressively decreasing combustion efficiency and temperature has been termed as “moderately absorptive (M-BrC)”, “weakly absorptive (w-BrC) and “very weakly absorptive (VW-BrC)”, respectively (Saleh et al., 2020; Moschos et al., 2024), as the k values decrease with lower burning temperature.

One of the most convenient and widely used methods for the characterization of optical properties of bulk BrC is the collection of aerosol particles on filters and subsequent extraction of the organic fractions with suitable solvents followed by filtration (Chen and Bond, 2010; Liu et al., 2013; Mo et al., 2017; Shetty et al., 2019; Li et al., 2020, Yan et al., 2020). The light absorption of solvent extracted OA is measured using a UV-vis spectrophotometer which provides high precision spectral data over a wide wavelength range. In previous studies water has been one of the primary solvents for extraction of BrC from filters (Bosch et al., 2014; Kirillova et al., 2014; Mukherjee et al., 2020) because of the atmospheric relevance of the water soluble organic (WSOC) fraction (Hallar et al., 2013, Taylor et al., 2017). Quantification of WSOC can be done with Total Organic Carbon (TOC) analysis (Li et al., 2016), and the loss estimation of the extraction process has very little uncertainty. On the other hand, the water insoluble BrC fraction is generally extracted from aerosol particles using methanol (MeOH) as a solvent (Chen and Bond, 2010) but it has been recently discovered that some highly light-absorbing, extremely low volatility organic (ELVOC) compounds may not be efficiently extracted by this procedure (Saleh et al., 2014; Liu et al., 2013). Quantifying the dissolved organic mass in MeOH is also challenging as organic solvents interfere with TOC measurements and indirect methods are used instead (Chen and Bond 2010; Cheng et al., 2017; Huang et al., 2018; Yan et al., 2020) to estimate methanol soluble organic carbon (MSOC) which leads to additional potential sources of uncertainties (Yan et al., 2020). Nevertheless, MeOH has exhibited very high



119 organic extraction efficiencies (Chen and Bond, 2010; Xie et al., 2017) and the optical properties of MSOC
120 have agreed well with OC extracted by more polar solvents like Dimethylformamide (DMF) for BB
121 emissions and coal combustion (Xu et al., 2022). Therefore, MeOH should be used in parallel to water to
122 extract BB emitted organic compounds with a broader range of polarities and gain more information on their
123 light-absorbing properties. Although filter based solvent extraction has its limitations, such as lack of
124 information on size-dependent absorption of extracted organics (Moosmüller et al., 2011; Liu et al., 2013;
125 Washenfelder et al., 2015), this analytical method is low-cost, easily accessible, and excludes the interference
126 of BC (or EC) and other light-absorbing species from OA absorption.

127 This work aims to define the emission factors of carbonaceous aerosols and characterize OA
128 optical properties for emissions originating from different open BB sources. We used a laboratory open
129 burning setup with the objective to create conditions representing natural Eurasian forest surface fires, in
130 which the combustion temperatures likely remain low and burning is dominated by smouldering (Rogers et
131 al., 2015; Walker et al., 2020). Finnish boreal peat and forest surface samples, commercially available boreal
132 peat samples, and permafrost peat from arctic Russia and Svalbard (Norway) were burned in the laboratory
133 setup. To assess the effects of burning conditions and biomass materials on carbonaceous emission factors
134 and their optical properties, we replicated conditions ranging from smouldering to flaming, with clearly
135 distinct combustion behaviour and MCEs for each fuel. Furthermore, we extended the study samples to
136 include South-African savanna biomass and North European wood stove emissions, thereby encompassing a
137 larger range of combustion conditions. Finally, we investigated the impact of atmospheric dilution and aging
138 on the chemical and optical properties of the organic aerosols by conducting environmental chamber
139 experiments either under photochemical or dark aging conditions for selected biomasses. The results were
140 used to derive imaginary refractive indices (k) for OAs, allowing for their classification within the black-
141 brown carbon continuum. The results are essential to accurately estimate the direct radiative forcing effects
142 of biomass burning emissions in climate models.

143

144 **2. Methodology:**

145

146 **2.1 Combustion setup and fuels**



147 The combustion experiments were conducted in the ILMARI laboratory of the Kuopio campus of University
148 of Eastern Finland (<https://sites.uef.fi/ilmari>) using an in-house designed open combustion appliance. The
149 appliance consisted of a steel cage, a concave plate, a metal/steel mesh and a metal/steel biomass holder. The
150 details of the combustion setup and the specific biomass holders used for each fuel type are illustrated in
151 Supplementary Fig. S2. Combustion was initiated using an electric resistor of which the power was adjusted
152 to generate exclusively ‘flaming’ or ‘smouldering’ emissions.

153 Eight different types of biomass were used as fuel samples in the combustion experiments. Namely,
154 commercially available peat samples (CP), Finnish boreal forest surface (BFS, including vegetation, litter,
155 and the soil organic layer), peat from two Finnish boreal peatlands (FIA and FIB), peat from arctic
156 permafrost regions of Russia (RUS) and Norway (NOR), savanna wood and grass from South Africa (SW
157 and SG respectively) were selected as fuels for this study. The origins and properties of the combusted
158 biomasses are available in the Supplementary Information (Section S1, Table S1). The savanna biomass
159 included in this study was part of the BASFAA (Boreal and savanna fire aerosol aging) measurement
160 campaign that took place from May to June 2022 at the ILMARI laboratory as described in Vakkari et al.
161 (2025). For this study specific combustion phases were selected for sampling representing either solely
162 flaming, pre-flame smouldering, or full combustion consisting of pre-flame smouldering, flaming and post-
163 flame smouldering periods. Therefore, a subset of Savanna biomass (SG and SW) combustions presented in
164 Vakkari et al. (2025), which included sampling of separate combustion phases (either flaming or pre or post
165 flame smouldering), were chosen to be included in this study. The sampling periods were selected based on
166 careful visual inspection of the combustion process and the concentrations of CO and CO₂, monitored online
167 as described below (Section 2.2).

168 **2.2. Sampling setup and gas phase measurements**

169 The emitted smoke was sampled using the setup presented in Fig. 1. We measured the gaseous

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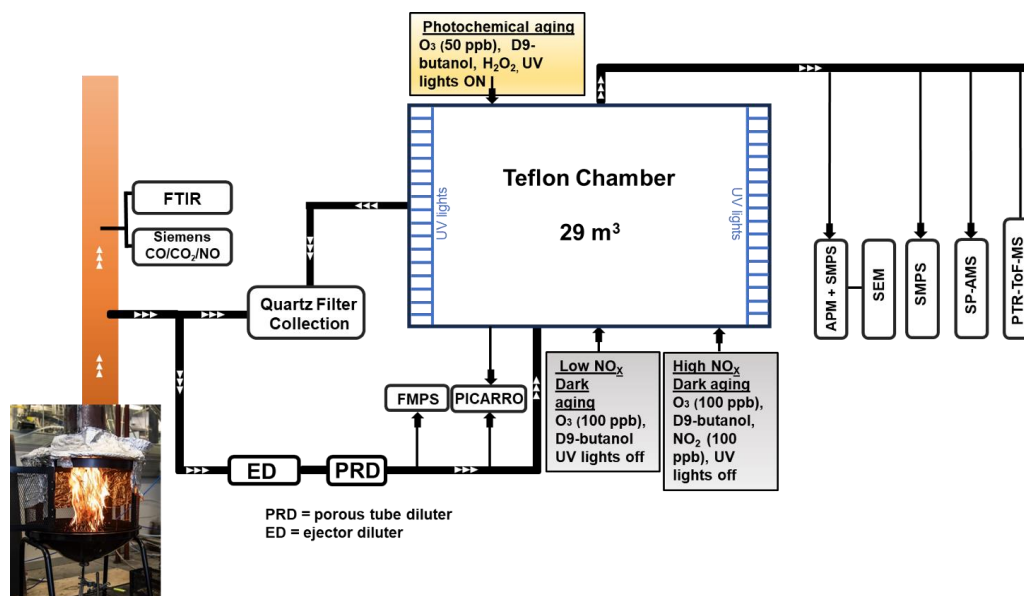


Figure 1: A schematic presentation of the experimental setup for combustion experiments involving the environmental chamber

compounds from the fresh, undiluted BB emission using an online multicomponent FTIR analyzer (Gaset Technologies Inc.), and the measured compounds are listed in Table S2. The fresh BB emissions were sampled through a PM₁₀ pre-cyclone and a heated probe (180 °C) before being diluted using an Ejector Diluter (Titta et al., 2016) and porous tube diluter (PRD). A dilution ratio (DR) of 4-20 was achieved during this two-stage dilution with clean synthetic air (Woikoski N50) in ambient temperature. A gas concentration analyzer (Picarro G2401) measured the concentrations of CO₂, CO, CH₄ and water vapor (H₂O) from the diluted exhaust (Fig. 1).

MCEs were calculated from the average increase in CO₂ and CO concentrations during the sampled combustion period relative to the background concentrations (Eq. 1)

$$MCE = \frac{\Delta[CO_2]}{\Delta[CO_2] + \Delta[CO]} \quad (1)$$

MCE was estimated from CO and CO₂ concentrations measured in both undiluted and diluted BB emission. Burns with average MCE values smaller than 0.9 were defined as smouldering-dominated combustion, while



187 the combustions with average MCE values larger than 0.9 were classified as flaming dominated (Table 1).

188

189 **2.3 Environmental chamber and aging of the emissions**

190 Different burning phases of Finnish BFS, and Finnish commercial peat (CP) samples were used in the
191 chamber experiments to study the effects of dilution and oxidative aging on the organic aerosol optical
192 properties. Additionally, samples of Savanna burning chamber experiments, described by Vakkari et al.
193 (2025), were analyzed as part of this study. For peatland samples FIA, FIB, NOR and RUS, only fresh
194 emissions were characterized in this study and no chamber feeding was performed. In each of the chamber
195 experiments, diluted primary BB emissions were sampled into a 29 m³ TeflonTM chamber (Leskinen et al.,
196 2015) which was pre-filled with purified air (Model 737-250, Aadco Instruments Inc.). For each experiment,
197 the sampling period to the environmental chamber was set based on the targeted mass concentration of 20-50
198 µgm⁻³ as estimated from the fresh particle size distribution measured online by a fast mobility particle sizer
199 (FMPS, model 3091, TSI Inc.). The relative humidity (RH) inside the chamber was maintained at 20% for
200 SG and SW emissions and at around 50% for BFS and CP emissions to reproduce typical daytime RH in
201 corresponding environments during the fire active seasons. The temperature in the chamber was kept
202 constant at approximately 22°C. The sample was first allowed to mix and homogenise for 20 minutes, which
203 was an adequate duration for the mass concentrations of particulate matter (PM) and organics inside the
204 chamber to stabilize, as observed respectively by a scanning mobility particle sizer (SMPS) and a soot
205 particle aerosol mass spectrometer (SP-AMS, Aerodyne Research Inc, USA) connected to the chamber. After
206 sampling of the fresh exhaust, Picarro was switched to monitor the gaseous components inside the chamber.
207 Total DRs inside the chamber were 135-4045 compared to the two-stage diluted fresh emission (Table 1),
208 based on the ratio of CO concentration measured by Picarro in the diluted exhaust and from the chamber.

209 After the stabilization period, chamber diluted primary BB emissions were monitored for
210 additional 45 mins before adding reactants to the chamber to induce oxidative aging. The reactants were
211 allowed to interact with the fresh BB emission in the chamber for 4.5 hours. For each experiment, 2 µL of
212 d9-butanol (~25 ppb) was injected into the chamber and the concentration was probed throughout the
213 experimental duration using a proton-transfer-reaction time-of-flight mass spectrometer (PTR-ToF-MS).



Four distinct aging conditions were simulated in the environmental chamber to evaluate the impact of aging on the physio-chemical and optical properties of different BB emissions. Firstly, photochemical aging conditions were induced inside the chamber for both flaming and smouldering combustion. In the photochemical experiments, UV-light in presence of externally fed O_3 (50 ppb) and H_2O_2 (0.5 ml of 30% v/v solution) led to the formation of hydroxyl radicals ($OH\cdot$). The OH exposure and equivalent photochemical age were determined from the decay of d9-butanol (Barnet et al., 2012). In our experiments, equivalent photochemical age ranged between 1-1.6 days (Table 1) with an assumed ambient OH concentration of $1.5 \times 10^6 \text{ cm}^{-3}$. In addition, two different dark aging conditions were simulated for two distinct sets of biomasses. The primary emissions from SG and SW underwent dark aging in the presence of externally added 100 ppb of O_3 and no additional NO_x . We classified this aging condition as “*low- NO_x dark aging*”, as the nitrate radical (NO_3) formation was limited by the lack of NO_x . CP and BFS burning emissions were aged in relatively ‘*high- NO_x* ’ conditions in the dark chamber with 100 ppb O_3 and 100 ppb NO_2 added externally. It should be clarified that even though we termed this aging condition as “*high- NO_x* ”, the NMVOC/ NO_x ratio was still relatively high (Seinfeld and Pandis, 2006) in the range of 4-6 in these cases (Table 1). Furthermore, seven experiments were conducted without any added oxidants in the chamber to evaluate the impacts of chamber dilution and other chamber processes on the exhaust emissions. Out of these seven experiments three were conducted by sampling emissions from the whole combustion of BFS, while the other four experiments were constituted with distinct smouldering and flaming dominated burns of CP and BFS.

2.4 Particle size distribution, density and morphology:

The particle size distributions were measured from the environmental chamber for the whole duration of each experiment by a scanning mobility particle sizer (SMPS) consisting of a differential mobility analyzer (DMA, model 3080; TSI Inc.) and a condensation particle counter (CPC, model 3775; TSI Inc.). In addition, an aerosol particle mass analyzer (APM, model 3602; Kanomax Inc.) in tandem with another SMPS (with DMA 3081 and CPC 3750, TSI Inc.) was used to measure the size distribution of mass classified particles to estimate the size-resolved effective densities of the particles (Leskinen et al., 2023; Mukherjee et al., 2024). Particle densities of primary emission from the chamber were measured before the application of any



241 additional reactants to the chamber. The density of the aged particles in the chamber were measured 4.5
242 hours after the addition and stabilization of the reactants (for dark aging) and/or turning on of the UV lights
243 (for photochemical aging; Table S4). Simultaneous to the density measurements, particles were collected on
244 holey-carbon grids (S147-4 Holey carbon film 400 Mesh Cu; Agar Scientific Inc.) from the chamber using
245 an aspiration sampler at a flowrate of 0.3 litres per minute (lpm). Subsequently, we performed Scanning
246 Electron Microscopy (SEM, Sigma HD/VP; Carl Zeiss NTS) to investigate the morphology of both chamber
247 diluted primary and aged BB emitted particles.

248 **2.5 Offline optical analyses:**

249 Biomass burning emissions were collected on precombusted 90 mm Quartz microfiber filter (Pallflex™
250 Tissuquartz™ 7203, Pall Corporation) from the fresh emissions and at the end of each chamber experiment.
251 First, the freshly emitted particles from the raw exhaust without any additional dilution were deposited on the
252 filter at a flow rate of 90 lpm. For experiments with subsequent chamber study, the sampling was done for
253 the same duration as the chamber feed (Fig.1). Approximately 4.5 hours after the chamber feeding, another
254 filter sample was collected from the chamber with the same flow rate for 120 minutes. For the boreal (FIA
255 and FIB) and arctic peat (RUS and NOR) samples, the primary emission was deposited on the filter
256 throughout the whole combustion period, and no chamber studies were performed.

257

258 Thermal-optical analyses were conducted on 1.5 cm² punches of the of the Quartz fiber filters (QMA)
259 containing deposited aerosol particles with the IMPROVE-A protocol (Chow et al., 2007) in an OC-EC
260 analyzer (Lab OC-EC Aerosol Analyzer; Sunset Laboratory Inc.). Additionally, two separate filter punches
261 from each experiment were extracted, one with ultrapure Mili-Q water (>18.2 MΩ) and the other with
262 methanol (Fisher Scientific, Analytical Reagent Grade ≥99.9% pure). After solvent extraction, the two
263 filters were dried under gentle airflow in a clean room for 12 hours before being analysed with the same OC-
264 EC analyzer. This setup (Fig. S4) enabled us to measure the dissolved organic concentration in the solvent
265 using Eq. 4 (see section 2.8) and therefore estimate the MAC_{OC} (Eq. 3).

266



267 **Table 1: List of chamber experiments with detailed parameterization of combustion and aging conditions (mean $\pm$ standard deviation)**

No.	Fuel	Combustion Condition (no. of replicates)	MCE	OC EF (fresh emission) (g kg ⁻¹)	EC EF (fresh emission) (g kg ⁻¹)	Final DR (chamber)	SMPS PM ₁ EF (g kg ⁻¹) (chamber primary)	GMD (nm)	Aging condition (in chamber)	NM VOC/NO _x (ppbC/ppb)	OH exposure (x 10 ¹¹ molecules.cm ⁻³ . s ⁻¹)	eqv. age (days)
1a	CP	flaming (n = 2)	0.971 $\pm$ 0.004	5.34 $\pm$ 2.32	0.29 $\pm$ 0.08	247 $\pm$ 106	6.98 $\pm$ 4.01	52.4 $\pm$ 1.60	photo aged	21.4 $\pm$ 4.40	2.09 ($\pm$ 0.33)	1.61 $\pm$ 0.26
1b		flaming (n = 1)	0.979	1.85	0.40	229	3.86	48.0	high NO _x dark aged	5.45	NA	
1c		flaming (n = 1)	0.986	1.39	1.81	273	8.05	54.4	no aging	6.40	NA	
1d		smouldering (n = 2)	0.619 $\pm$ 0.147	58.3 $\pm$ 31.5	0.24 $\pm$ 0.07	764 $\pm$ 118	152 $\pm$ 89	48.3 $\pm$ 5.20	photo aged	88.6 $\pm$ 5.01	1.15 ($\pm$ 0.63)	0.88 $\pm$ 0.48
1e		smouldering (n = 1)	0.793	26.8	0.16	429	33.6	45.5	high NO _x dark aged	4.20	NA	
1f		smouldering (n = 1)	0.479	73.7	0.49	666	228	59.7	no aging	108	NA	
2a	BFS	flaming (n = 3)	0.922 $\pm$ 0.003	3.20 $\pm$ 0.94	0.43 $\pm$ 0.15	309 $\pm$ 199	6.07 $\pm$ 1.05	51.0 $\pm$ 4.10	photo aged	63.9 $\pm$ 41.7	1.80 ($\pm$ 0.34)	1.39 $\pm$ 0.26
2b		flaming (n = 3)	0.952 $\pm$ 0.005	3.18 $\pm$ 0.70	0.39 $\pm$ 0.08	387 $\pm$ 140	4.79 $\pm$ 0.82	51.6 $\pm$ 0.70	high NO _x dark aged	4.78 $\pm$ 1.23	NA	
2c		flaming (n = 1)	0.961	3.06	0.73	833	5.95	53.6	no aging	20.0	NA	
2d		smouldering (n = 3)	0.759 $\pm$ 0.020	23.9 $\pm$ 15.7	0.08 $\pm$ 0.07	276 $\pm$ 100	48.3 $\pm$ 29.5	84.7 $\pm$ 41.4	photo aged	133 $\pm$ 104	1.57 ($\pm$ 0.11)	1.21 $\pm$ 0.09
2e		smouldering (n = 3)	0.761 $\pm$ 0.017	24.0 $\pm$ 10.5	0.07 $\pm$ 0.06	287 $\pm$ 123	26.6 $\pm$ 21.1	73.2 $\pm$ 22.8	high NO _x dark aged	23.5 $\pm$ 19.6	NA	
2f		smouldering (n = 1)	0.774	15.3	0.33	135	23.96	51.2	no aging	163	NA	
3a	SW	flaming (n = 1)	0.935	8.04	0.97	922	12.4	60.5	photo aged	26.1	1.41	1.09
3b		flaming (n = 1)	0.950	5.93	0.60	244	10.5	48.0	low NO _x dark aged	43.9	NA	
3c		smouldering (n = 1)	0.589	87.0	5.18	389	61.1	81.0	photo aged	53.2	1.29	0.99
3d		smouldering (n = 1)	0.892	22.5	1.36	416	31.1	62.6	low NO _x dark aged	69.7	NA	
4a	SG	flaming (n = 1)	0.943	4.83	0.31	512	6.24	48.1	photo aged	124	1.46	1.13
4b		flaming (n = 1)	0.973	2.87	0.28	1974	5.44	50.3	low NO _x dark aged	2.73	NA	
4c		smouldering (n = 1)	0.803	31.8	1.19	4045	28.6	60.5	photo aged	19.3	1.49	1.15
4d		smouldering (n = 1)	0.843	14.0	1.25	902	23.2	58.5	low NO _x dark aged	108	NA	



269 For filters containing fresh emissions collected from the raw exhaust, 40 mL of solvent was used for the
270 extraction of the 1.5 cm² filter area. The volume of solvent used for extraction of highly loaded filters
271 containing fresh emissions can be a limiting factor for dissolved OC concentration, especially for BB emitted
272 particles that have been shown to be dominated by non-polar molecules (Lin et al., 2017) that are less soluble
273 in water. We chose 40 ml of solvent, as this volume provided a dilute micromolar solution in which the light
274 absorption by the dissolved organics would be linearly proportional to their concentration, in accordance
275 with Beer-Lambert's law (Huang et al., 2018). The utilized solvent volume was also sufficient to dissolve the
276 sparingly soluble low volatility strongly light-absorbing fraction of OC. On the other hand, for the chamber
277 diluted fresh emissions and aged particles, 3 cm² filter punches were extracted in 20 mL water or methanol
278 (Cao et al., 2021; Fan et al., 2018) to obtain solution phase organic concentrations necessary for analytically
279 significant S/N ratio in the UV-vis spectrophotometry.

280 In parallel to the deposition of primary particles from the raw exhaust on quartz fiber (QMA) filters for
281 emission factor calculation, a fraction of the emitted particles was fed to the TeflonTM chamber to create real
282 life atmospheric dilution. It is important to study these particles in a diluted system to understand their
283 chemical and optical evolution in the atmosphere. The particle size distribution, effective density, and
284 morphology of the primary BB emission were studied after feeding them in the chamber and letting them
285 mix homogeneously with clean air. Subsequently, different oxidants were added to the chamber to initiate
286 photochemical or dark oxidation reactions. After 4.5 hours of feeding the oxidants, the particles from the
287 chamber were deposited on another QMA filter. In addition to that, five separate experiments were conducted
288 for different BB emissions, where no oxidants were added externally to the chamber and the primary
289 particles were on QMA filter 4.5 hours after the chamber feed. These samples are referred to as "chamber
290 primary" samples. The QMA filters containing chamber aged and chamber primary particles were produced
291 to undergo OC-EC analyses and solvent extraction similar to the QMA particles collecting raw exhaust
292 primary particles.

293 The filter punches were submerged in the solvents and sonicated for 10 mins in three separate intervals in an
294 optimized ultrasonicator (SONOREX Digitec, Bandelin Inc.). The optimization enabled a much gentler
295 sonication as compared to regular commercial ultrasonicators (Huang et al., 2018; Li et al., 2020), resulting



296 in high extraction efficiencies without damaging the filter punch and thereby allowing subsequent OC-EC
297 analyses. Such gentle ultrasonication also minimised the dislodging of insoluble organics or EC from the
298 filter to the solution (Phillips and Smith, 2017). During the intervals between the 10-minute sonication
299 windows, the samples were kept in a refrigerator to keep the effects of any thermal disintegration of OC or
300 chemical transformation through reactions with the solvent (Bateman et al., 2008; Chen et al., 2022) due to
301 the added kinetic and thermal energy during the sonication at minimal.

302 After the sonication, the elute was passed through 0.2 µm hydrophilic PTFE syringe filters (Fisherbrand) to
303 remove any insoluble particles from the solution. Aliquots of 3 mL from the filtered solutions were taken in a
304 quartz cuvette with 1 cm path length and UV-vis spectra were recorded for the wavelength range of 250-700
305 nm using a spectrophotometer (UV-2401 PC, Shimadzu). The overall contact time between the solvent and
306 the extracted organics in the solution never exceeded 2 hours before the recording of the optical spectra.
307 These were much smaller time intervals compared to the reaction rates between methanol and carboxyl
308 groups for instance (McIntyre and McRae, 2005), enabling us to disregard any potential artifact caused by
309 the solvent (methanol) to the chemical composition of the original OC deposited on the filter (Lin et al.,
310 2012).

311 **2.6. Residential wood combustion experiments**

312 In addition to different open BB emissions, filter samples from experiments performed with modern
313 European logwood-fired chimney stove (Aduro 9-3) operated with beech logs from two separate
314 experimental campaigns were included in the study. This enabled us to extend the analysed sample set to
315 conditions representing high-temperature biomass combustion, and to test the validity of the BrC-BC
316 continuum. The general experimental setups of these two sets of experiments have been presented in our
317 previous works (Ihalainen et al., 2019a, Mukherjee et al., 2024). Each experiment consisted of four
318 (Ihalainen et al., 2019a) or six (Mukherjee et al., 2024) 45 min batches of wood log burning, ignited with
319 wood sticks as kindling. Photochemical aging of the residential wood combustion (RWC) exhaust was
320 conducted by the Photochemical Emission Aging flow tube reactor (PEAR, Ihalainen et al., 2019b), in
321 similar conditions as the low-dilution aging in Hartikainen et al. (2020). Briefly, exhaust samples were
322 collected onto quartz fiber filters before and after the PEAR in a flow rate of 10 lpm for 60 min, with each



sample consisting of one full batch and 20 min of the subsequent batch, either from the cold stove (consisting of 1st and 2nd batch of wood) or in the warm stove (consisting of 3rd and 4th batches of wood) combustion. The filter samples were extracted similarly to the open biomass burning samples for optical analyses.

2.7. Optical data processing:

The raw absorption (A) of the extracted solution measured by the UV-vis spectrophotometer was converted to the light absorption coefficient (β_{abs} , Mm^{-1}) in the sampled air using Eq. 2 (Hecobian et al., 2010):

$$\beta_{abs}(\lambda) = \frac{(A_{\lambda} - A_{700}) \times V_l \times \ln(10)}{V_a \times l} \quad (2)$$

where A_{λ} is the raw absorbance measured by the spectrophotometer in base-10 at wavelength λ and A_{700} is the absorption at 700nm (which was subtracted from A_{λ} to correct for baseline drift caused by the absorption), V_l is the volume of the extract, V_a (m^3) is the volume of the air sampled through the quartz filter during emission measurements, l is the path length of light through the solution, which is equivalent to the width of the quartz cuvette (0.01 m):

The mass absorption coefficient (MAC) of the extracted BrC at a particular wavelength λ is given by Eq. 3 (Liu et al., 2013):

$$MAC(\lambda) = \frac{\beta_{abs}(\lambda)}{C} \quad (3)$$

where C is the concentration of the dissolved organic in a particular solvent (g m^{-3}). C for methanol and water soluble organics (C_{MSOC} and C_{WSOC} respectively) were calculated as:

$$C_{MSOC} = \frac{(\text{OC in original filter}) - (\text{OC in methanol extracted filter})}{\text{volume of methanol used for extraction}} \quad (4a)$$

and,

$$C_{WSOC} = \frac{(\text{OC in original filter}) - (\text{OC in water extracted filter})}{\text{volume of water used for extraction}} \quad (4b)$$

C_{WSOC} was also directly measured from the aqueous filter extracts using a total organic carbon (TOC/TN) analyzer (TOC-L, Shimadzu) to compare with the estimates obtained using Eq. 4b. Overall the two methods



yielded comparable C_{WSOC} values ($R^2 = 0.98$, Fig. S5.), so for consistency we have used the filter based C values for both MSOC and WSOC in calculating MAC. The imaginary part of the refractive index, k , was related to MAC at any given wavelength λ according to Jennings et al. (1979):

$$MAC(\lambda) = \frac{4\pi k(\lambda)}{\rho \lambda} \quad (5)$$

where ρ is the density of the dissolved organic (g cm^{-3}). For fresh, photochemically aged and dark aged emissions different mean density values were used for calculating k as mentioned in Table S4. For the RWC experiments, particle densities measured for fresh and PEAR-aged beech combustion exhaust particles from the same appliance (Mukherjee et al., 2024). were applied for the calculation of k . Specifically, density values of $1(\pm 0.1) \text{ g cm}^{-3}$ and $1.6(\pm 0.1) \text{ g cm}^{-3}$ were used for the primary and aged RWC exhaust particles, respectively.

Absorption Ångström exponent (AAE) of MSOC and WSOC were estimated between a pair of wavelengths λ_1 and λ_2 as the power law exponent of the ratio between the absorption coefficients at the two wavelengths (Moosmüller et al., 2009), such as:

$$AAE(\lambda_1, \lambda_2) = \frac{\ln\{\beta_{abs}(\lambda_1)/\beta_{abs}(\lambda_2)\}}{\ln(\lambda_2/\lambda_1)} \quad (6)$$

For this study, AAE between wavelengths 300-550 nm were used to describe the optical behaviour of BrC emitted from different biomass sources. The wavelength dependence of the imaginary refractive index k is denoted as w (Saleh et al., 2014) which we derived for fresh and chamber samples using a similar power law relationship as Eq. 5 but involving k at two different wavelengths. In literature, w has also been estimated using AAE (Saleh et al., 2014, McClure et al., 2020) as they are related by:

$$AAE \approx w + 1 \quad (7)$$

2.8 Emission Factor Calculation

Emission factors from the combustion experiments were calculated by the carbon mass balance method such (Eq. 8; Yokelson et al., 1999)



$$EF_X = F_c \cdot 1000 \cdot \frac{ER_X}{\frac{\sum \Delta C}{\Delta CO}} \quad (8)$$

where F_c is carbon fraction in the combusted sample and the summation $\sum \Delta C$ is the total carbon released during the combustion. For savanna biomasses (SG and SW) $\sum \Delta C$ was estimated as (Vakkari et al., 2025):

$$\sum \Delta C = \Delta CO_2 + \Delta CO + \Delta CH_4 + \Delta OC + \Delta rBC + \sum \Delta C_{VOC} \quad (9 a),$$

ΔCO_2 , ΔCO and ΔCH_4 were their respective concentrations in ppm subtracted by background and synthetic air concentration, ΔOC was measured by AMS, rBC was measured by SP2 and VOCs were measured by VOCUS. For commercial peat (CP), boreal forest surface (BFS) and natural peatland (FIA, FIB, RUS and NOR) emissions, $\sum \Delta C$ was estimated as:

$$\sum \Delta C = \Delta CO_2 + \Delta CO + \Delta CH_4 + \Delta OC + \Delta EC + \sum \Delta C_{VOC} \quad (9 b),$$

For CP and BFS, CO_2 , CO and CH_4 were measured using Picarro and non-methane VOCs (NMVOC) were estimated from FTIR (Table S2). For natural peatland samples CO_2 , CO , CH_4 and NMVOCs were all measured by FTIR. OC and EC for all of these samples were estimated from $OC-EC$ analyses of filters.

ER_x in Eq. 4 is the ratio of measured concentration in $\mu g m^{-3}$ relative to the CO carbon concentration in $\mu g C m^{-3}$. ER_x was calculated by Eq. 5, where the CO concentration measured in ppm is converted to CO carbon concentration in $\mu g C m^{-3}$ using the following equation:

$$\Delta ER_X = \frac{X}{\frac{12.01 g/mol \cdot 101325 Pa}{8.31451 J / mol \cdot K \cdot (295.15 K + T)} \cdot \Delta CO} \quad (10)$$

Where X is the concentration of the measured parameter in $\mu g m^{-3}$, ΔCO is the CO concentration in the emission in ppm compared to background, T is measured temperature in the chamber in Kelvin (295.15 K or 22°C), 12.01 is the molecular weight of carbon ($g mol^{-1}$), 101325 standard atmospheric pressure in Pa, 8.31451 is gas constant in SI unit.

2.9 FT-ICR MS analyses:

Chemical compositions of the biomass burning emission samples collected on quartz filters were characterized by means of ultrahigh-resolution mass spectrometry. All experiments were performed using a



12-T Bruker solariX XR Fourier transform ion cyclotron resonance mass spectrometer (FT-ICR MS) (Bruker Daltonics GmbH, Bremen, Germany), equipped with a dynamically harmonized ICR cell (ParaCell®). The mass spectrometer was coupled to an Apollo-II atmospheric pressure chemical ionization (APCI) source, fitted with a direct insertion probe (DIP) accessory. This set-up enabled chemical characterization of the filter samples directly with minimal sample preparation. Five layers of each quartz filter sample were packed into a prebaked glass capillary (Hirschmann melting point tube) and inserted into the ion source vaporizer, held at 370 °C. The capillary voltage was set to 4500 V and corona current to 4000 nA. Dry nitrogen was used as the drying (3.5 Lmin⁻¹, 220 °C) and nebulizing (2.0 bar) gas. After an induction time of 10 s, the MS data were recorded until the total ion current plateaued (i.e number of scans ranged from 15 to 35), indicating that the entire aerosol sample was completely desorbed at the given temperature.

All DIP-APCI FT-ICR measurements were conducted in a positive ion mode. The generated ions were accumulated in the hexapole ion trap and transferred to the ICR cell for trapping, excitation and detection. The instrument control and data acquisition were performed by Bruker fimsControl 2.1 software. For each spectrum, a broadband frequency excitation and detection were carried out with 4 MWord time-domain transients (transient time 1.1 s), which were full-sine apodized and zero-filled once to provide the final 8 MWord magnitude-mode data spanning *m/z* range of 100–2000. The time-of-flight and ion accumulation time settings were 1.0 ms and 0.30 s, respectively.

The FT-ICR instrument was externally *m/z*-calibrated prior to the measurements using a polystyrene (PS) standard covering a wide mass range and reaching accuracies generally below 1 ppm. The initial spectral post-processing was done with Bruker DataAnalysis 5.1 software, including the internal re-calibration of the *m/z*-axes with a custom-made mass list for organic aerosol samples. For the peak picking, a signal-to-noise (S/N) ratio was set at ≥6. PetroOrg IS 18.0.3 software (Omics LLC, Tallahassee, FL, USA) was used for the molecular formula assignments with the following constraints: double bond equivalent (DBE) 0–40; mass error ±1.0 ppm; atomic formula ¹²C₁₋₁₀₀¹H₁₋₂₀₀¹⁴N₁₋₄¹⁶O₁₋₁₅³²S₁₋₄. The elemental composition boundaries for the annotation were chosen based on careful manual inspection of spectra identifying the edges of the observed chemical space. The time-resolved spectral information of the DIP experiment had been summed to an average spectral read back for each measurement.



421 For data interpretation and visualization of the complex lists of attributed sum formulae, we used established
422 data grouping and fingerprint diagrams (Schneider et al., 2024a; 2024b). Visualization and pre-processing,
423 calculating molecular properties and diagnostic measures from the sum formulae were performed via
424 MATLAB (MATLAB R2023a, MathWorks Inc., MA). Characteristic molecular properties encompassed
425 double bond equivalents (DBE), aromaticity index (AI), saturation vapor pressure (C^*) and average carbon
426 oxidation state (OS_c) frequently used in ultra-high resolution mass spectrometric studies of complex
427 environmental sample materials (Koch and Dittmar, 2006; Kroll et al., 2011; Li et al., 2016)

428 **2.10: SP-AMS analyses:**

429 This study employed a soot particle aerosol mass spectrometer (SP-AMS; Aerodyne Research Inc., Billerica,
430 MA, USA; Onasch et al., 2012) to analyze the concentration levels, mass spectral signatures, and size
431 distributions of non-refractory (organics, sulfate, nitrate, ammonium, and chloride) and refractory (e.g.
432 metals, rBC) components. The SP-AMS enhances the capabilities of the standard AMS by incorporating a
433 laser vaporizer, which enables the analysis of refractory aerosol components. While the instrument can
434 function using only the laser vaporizer, both the laser and tungsten vaporizers were used in this study to
435 ensure comprehensive detection of mentioned chemical species. Size-resolved measurements were obtained
436 using particle time-of-flight (PToF) mode, with the SP-AMS aerodynamic lens enabling detection of
437 particles ranging from roughly 50 nm to 1 μ m. The instrument operated with a time resolution of 60 seconds,
438 with about half of the time measuring in mass spectrum mode and the other half in PToF mode. A calibration
439 of the SP-AMS based on particle mass was carried out using size-selected, dried particles of ammonium
440 nitrate and ammonium sulfate. This approach allowed for the determination of an effective nitrate response
441 factor, as well as the relative ionization efficiencies (RIEs) for ammonium (RI_{NH_4}) and sulfate (RI_{SO_4}),
442 by converting the instrument signals into nitrate-equivalent mass concentrations. The determined RIE value
443 for NH_4 was 3.4 while the one for sulfate was 0.9. A default RIE value of 1.4 was used for organics. SP-
444 AMS data processing was conducted using the AMS analysis tools SQUIRREL (version 1.63B) and PIKA
445 (version 1.23B) within Igor Pro 8 (Wavemetrics, Lake Oswego, OR).

446



3. Results

3.1 Emission factors of OC and EC in fresh emission

The emission factors of OC (EF_{OC}) and EC (EF_{EC}) were determined based on thermal-optical carbon analyses of filters collected from fresh exhaust aerosol (Fig. 2, Fig. S6). Generally, we observed higher EF_{OC} for smouldering combustions with low MCE, as expected. For EC however, the trend was reversed, with flaming combustion having higher MCE producing more EC per kg of fuel (Fig. S6). The combustion emissions in our experiments were OC rich, most likely due to the open setup of the burner, which resulted in lower combustion temperature compared to wood stove emissions, due to high air-to-fuel ratios and because there is no combustion chamber around the fire to retain the heat in the surroundings. However, there were significant variations in EF_{OC} obtained from different combustion conditions and for different biomasses, ranging from 1.39-7.44 g kg⁻¹ for flaming dominated combustions and 6.12-89.9 g kg⁻¹ for smouldering dominated combustions. CP and SW samples had the largest EF_{OC} of 53.4(±29.0) g kg⁻¹ and 50.6(±29.8) g kg⁻¹ respectively, during the smouldering combustion phase. In contrast, full combustion of different natural peat samples yielded EF_{OC} in the range of 5.06-32.7 g kg⁻¹, with FIA (at depth of 30-60 cm) having the smallest (6.09±1.2 g kg⁻¹) and FIB (at depth of 30-60 cm) having the largest (29.8±2.92 g kg⁻¹) average EF_{OC} among peat samples. Interestingly, the average MCE values and EF_{OC} reported for field measurements of subtropical Indonesian peat fires (Stockwell et al., 2015; Jayarathne et al., 2018) and laboratory studies of North American (Black et al., 2016; Chakrabarty et al., 2016), West European (Iinuma et al., 2007) and South-east Asian peat fires (Christian et al., 2003; Iinuma et al., 2007) fell inside the range of our reported EF_{OC} for laboratory combustion of Finnish (boreal) and arctic peat samples. In our study, flaming and smouldering combustion of BFS samples had average EF_{OC} of 3.17(±0.77) g kg⁻¹ and 21.5(±13.8) g kg⁻¹ respectively, while previously reported EF_{OC} for North American boreal forest fires were in similar range of 5.9(±2.5) g kg⁻¹ (Andreae, 2019). Savanna wood and grass (SW and SG) samples burnt under flaming conditions in this study were estimated to have average EF_{OC} of 6.47(±0.98) g kg⁻¹ and 3.56(±0.91) g kg⁻¹ respectively, which were somewhat higher in comparison to previously reported organic emission factor of 3.0(±1.5) g kg⁻¹ (Andreae, 2019) and 2.62(±1.24) g kg⁻¹ (Akagi et al., 2011) from high MCE savanna fires.



474 Emissions from savanna biomass had higher average EF_{EC} compared to the other biomasses,
475 with an average EF_{EC} of $3.03(\pm 1.77) \text{ g kg}^{-1}$ and $0.94(\pm 0.16) \text{ g kg}^{-1}$ from the smouldering emissions of the
476 woody (SW) and grassy (SG) fuels respectively. The carbonaceous fraction of flaming dominated emissions
477 from the woody savanna samples ($MCE \sim 0.94 \pm 0.01$) consisted of $12(\pm 2) \%$ of EC while savanna grass had
478 an EC content of $7.4(\pm 1.5)\%$ ($MCE \sim 0.96 \pm 0.01$). Previously reported average EF_{EC} values for savanna and
479 grassland fires ranged from $0.37(\pm 0.20) \text{ g kg}^{-1}$ (Akagi et al., 2011) to $0.53(\pm 0.35) \text{ g kg}^{-1}$ (Andreae, 2019)
480 while Vakkari et al. (2018) reported EF_{EC} of 0.67 g kg^{-1} for South African savanna grass burning. These
481 values fell in between our estimated EF_{EC} from flaming burn of savanna grass ($0.27 \pm 0.01 \text{ g kg}^{-1}$) and wood
482 ($0.72 \pm 0.17 \text{ g kg}^{-1}$). Full combustion ($MCE 0.74-0.89$) of the Finnish peat samples (FIA and FIB) collected at
483 the depth of 30-60 cm from the surface level and arctic permafrost peat samples from Svalbard (NOR) and
484 Russia (RUS) had the lowest measured average EF_{EC} (in the range of $0.02-0.13 \text{ g kg}^{-1}$), which matched well
485 with previously reported EF_{EC} from field and lab studies of subtropical Indonesian (Stockwell et al., 2015;
486 Christian et al., 2003) and laboratory replicates of North American peat fires (Black et al., 2016). Flaming
487 and smouldering combustions of CP had average EF_{EC} of $0.70(\pm 0.65) \text{ g kg}^{-1}$ and $0.28(\pm 0.14)$ which
488 resembled EF_{EC} reported from laboratory studies of Western European and South-East Asian peat fires
489 (Iinuma et al., 2007). Burning of the boreal forest surface sample (BFS) resulted in higher EC emissions than
490 any of the peat samples with an average EF_{EC} of $0.45(\pm 0.16) \text{ g kg}^{-1}$ for high MCE ($\sim 0.94 \pm 0.01$) combustion
491 and an average EF_{EC} of $0.11(\pm 0.09) \text{ g kg}^{-1}$ for low MCE ($\sim 0.76 \pm 0.02$) smouldering burns. For North
492 American boreal forest fires, Andreae (2019) reported EF_{EC} of $0.43(\pm 0.21) \text{ g kg}^{-1}$ with the fire having an
493 average MCE of $0.89(\pm 0.04)$, which resembles Finnish BFS EF_{EC} for flaming dominated burns. Thus, it
494 seems that our experiments with simulated surface fire conditions at comparably lower combustion
495 temperature typical for Eurasian wildfires yield similar EC emissions as reported for North American
496 wildfires that are commonly high temperature crown fires.

497 In comparison, for residential wood combustion (RWC) of North European beech wood logs
498 (Mukherjee et al., 2024) the respective average OC and EC emission factors were $0.64(\pm 0.31) \text{ g kg}^{-1}$ and
499 $0.28(\pm 0.04) \text{ g kg}^{-1}$, respectively. Therefore, the OC emission factors for the open burning in the current study
500 were approximately 4-140 times higher for smouldering emissions and 2-13 times higher for flaming



emissions than our previous estimates of RWC emission. In comparison, we did not observe such a significant difference between the EC emission factors of RWC and flaming BB emissions. This implies that surface fires have the potential to be important sources of specifically organic pollutants and BrC in forest fire prone environments, whereas the BC emissions would not be particularly strong from these fires.

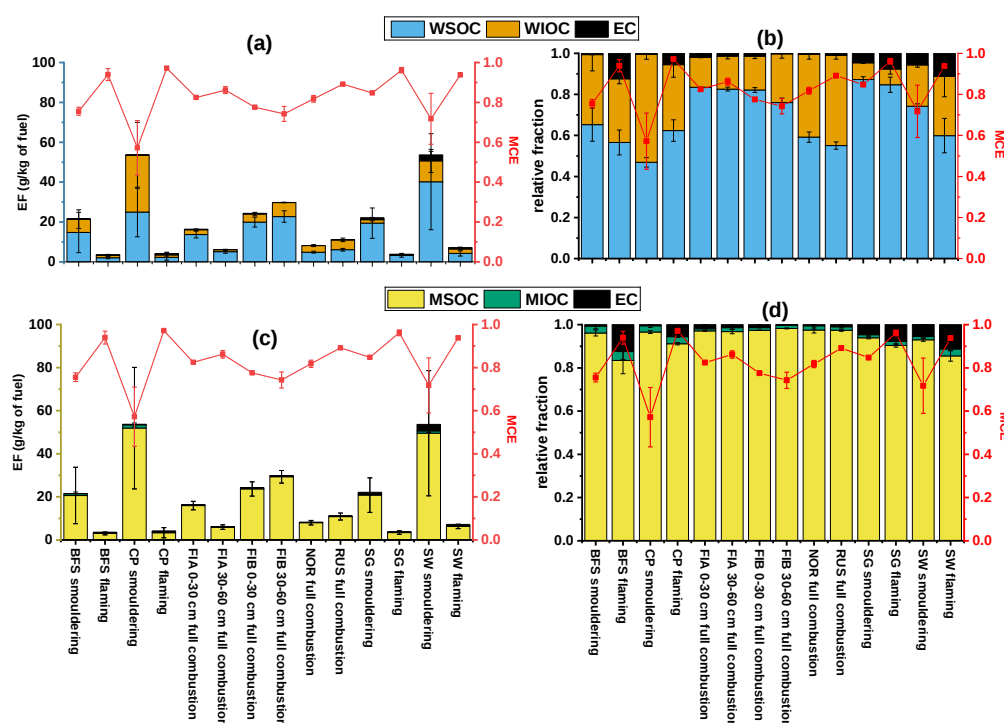


Figure 2. Emission factors of water-soluble (WSOC), methanol-soluble (MSOC), water-insoluble (WIOC), methanol-insoluble organic carbon (MIOC) and elemental carbon (EC) in absolute scale (a,c) and relative to total EF (b,d).

3.2 Chemical composition and solubility of fresh organic aerosol

Chemical characterization of the fresh emission samples from whole combustions of peat (FIA, FIB, RUS and NOR) and Finnish BFS have been previously reported by Schneider et al. (2024a) in which electrospray ionization (ESI) was utilized for the FT-ICR mass spectrometry on extracts, targeting the polar constituents. For this study, the filters containing specifically flaming and smouldering emissions of CP and BFS were analyzed directly from the filters by using 12-T DIP-APCI FTICR-MS to identify differences in the organic



515 matter compositions due to both fuel and combustion conditions. The advantage of the DIP-APCI technique
516 used in this study is the ability to detect analytes with low solvent extraction recoveries due to minimal sample
517 preparation required (Rüger et al., 2021). Additionally, APCI was able to address a wide chemical space
518 covering polar up to apolar analytes. Therefore, inspite of the lower mass loadings on the filters for specific
519 combustion phases compared to full biomass combustions as reported in Schneider et al. (2024a), we could
520 assign up to 2000-4000 monoisotopic elemental compositions. Overall, CHO, CHNO and CH compound
521 classes were the most abundant in the fresh BB emissions of CP and BFS. The low temperature smouldering
522 combustions have been shown to emit mostly direct thermal degradation products from biomass pyrolysis
523 (Chakrabarty et al., 2016) compared to flaming combustion (Engling et al., 2006, Popovicheva et al., 2019),
524 with similar chemical composition and molecular structures to that of the fuel (Kourtchev et al., 2011). Similar
525 to the smouldering dominated peat burning emissions reported in Schneider et al. (2024a), we also observed
526 that the CP smouldering emissions were chemically most diverse with the highest number of sum formulae
527 unique to this burning condition (in total 774), thus reflecting the complex inherent compositional variability
528 of the peat samples (Figure 3). CP smouldering emissions also had clearly higher CHN and CHNO fractions
529 than all the other biomasses and combustion conditions. This observation can be attributed to the fact that the
530 CP samples contained only below-surface peat that have decomposed for longer time periods and experienced
531 more microbial activity than the BFS samples, leading to higher N containing species. Flaming combustion of
532 CP emitted the second highest number of unique sum formulae (in total 297) but had much lower CHN class
533 content compared to smouldering conditions, suggesting a considerable effect of the burning condition / MCE
534 on the POA chemical composition (Figure 3).

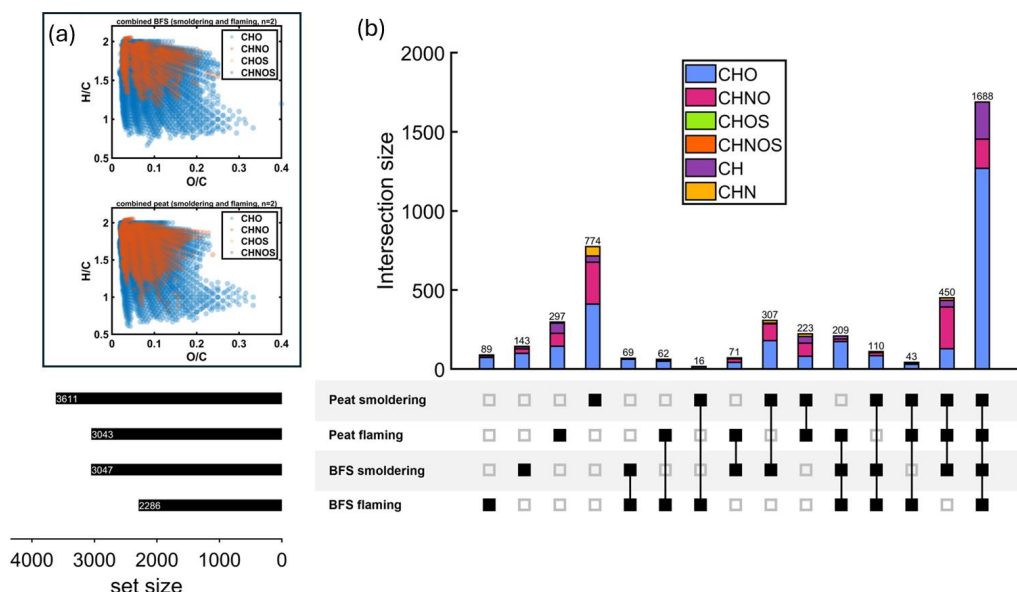


Figure 3: (a) Van-Krevelen diagrams of oxygen-containing compounds and (b) upset plot intersections (combined CP and BFS) with compound class indicated by colour. Samples were analyzed by positive-ion DIP-APCI FT-ICR mass spectrometry. Upset plot of four organic aerosol datasets with the number of measured molecular formulae in the whole dataset (bottom left) and the individual intersections (top right) indicated by color (CHO-light blue, CHOS-green, CH-purple, CHNO-magenta, CHNOS-orange and CHN-yellow)

On the other hand, the BFS smoldering and flaming emissions exhibited chemically similar compounds, with nearly 1900 identified molecular formulae shared among all the analyzed BFS samples. Almost none of the molecular formulae were unique to BFS smoldering or flaming samples, as the variability among replicates themselves was high. This indicates that although different combustion conditions of BFS strongly influence the OC emission factor, it does not lead to prominent chemical differences in POA. Similar to CP, BFS emissions were also dominated by CHO, CHNO and CH classes, although a much lower fraction of CHN was present in the BFS smoldering emission compared to CP. Furthermore, a more detailed inspection of the identified formulae of the nitrogen-containing molecules in the fresh CP and BFS emissions reveals that they mostly belonged to the CHNO class, contained one nitrogen substitution, and consisted of both aliphatic and aromatic nitro-organic compounds (Figure 4). Specifically, nitroaromatics were more



abundant in the peat samples, which is relevant, as nitroaromatics are known to be prominent BrC chromophores absorbing light in the wavelength range of 360-600 nm (Fleming et al., 2020).

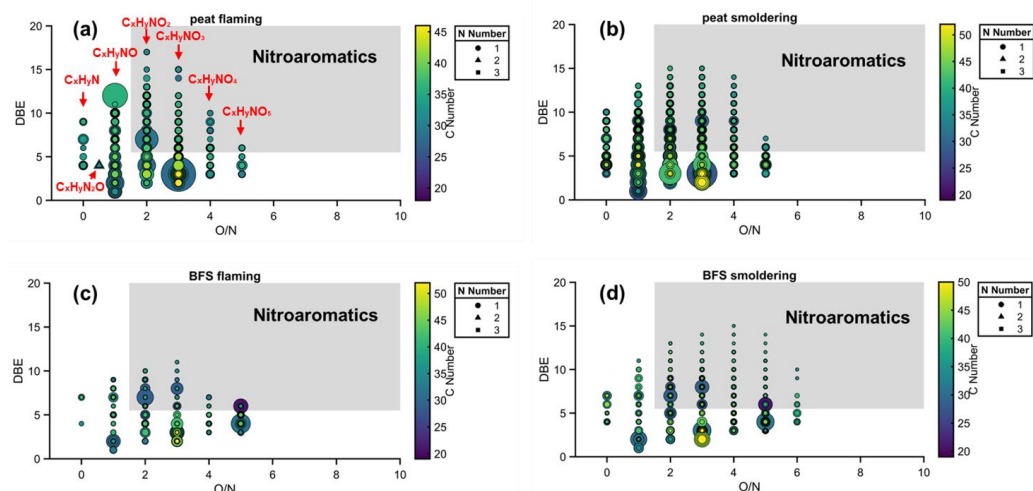


Figure 4: Detailed chemical characterization of CHN and CHNO compounds in (a) CP flaming, (b) CP smouldering, (c) BFS flaming and (d) BFS smouldering fresh emissions. The x-axis shows O:N ratio and y-axis corresponds to double-bond equivalence (DBE). Nitroaromatic compounds fall into the shaded (grey) area in the figures.

Due to their different chemical compositions, OC from various fresh BB emissions exhibited variable solubility in water and methanol (Fig. 2, Fig. S11). While methanol extracted almost all the OC (~95%) from the filter irrespective of fuel or combustion type, the water-soluble fraction of OC (WSOC) exhibited fuel dependent behaviour. For example, both FIA and FIB emissions exhibited higher WSOC fractions (76-83%) in the total organic carbon (TOC) compared to arctic permafrost peat samples NOR (~59% of TOC) and RUS (~55% of TOC). This has been previously elaborated by Schneider et al. (2024a), who found that in comparison to FIA and FIB samples, NOR and RUS peat combustion emissions were constituted by less oxidized primary organic carbon (POC), which explains the diminished solubility in a polar solvent like water (Budisulistiorini et al., 2017). Among the biomass samples used in this study, SG showed the highest WSOC fractions for both flaming (~92% of TOC) and smouldering (~91% of TOC) combustion (Fig. 2c-d), suggesting abundance of more oxidized POC in these emissions, while smouldering burning of CP consisted of the smallest WSOC fraction (~46%). These observations point towards the presence of non-polar organic



moieties, which are soluble in methanol but not in water, as previously reported in the literature (Lin et al., 2017).

3.3 Optical properties of Fresh organic aerosol:

Due to the variation in chemical composition and solubility of the fresh BB emissions, we observed substantial diversity in the optical properties of the freshly emitted organic particles. Since MSOC corresponded to around 95% of the total OC in our fresh BB emissions, we can use MSOC as a proxy for BBOA in our study. The absorption angstrom exponent ($AAE_{300-550}$, measured in the range of 300-550 nm) of the particles sampled from the fresh emission for different fuels ranged between 4.4-5.3 and 4.7-6.2 for WSOC and MSOC, respectively. These values fall in the ranges of previously reported AAE for BB emitted as well as urban WSOC and MSOC (Cao et al., 2021; Fan et al., 2018; Yan et al., 2015; Mukherjee et al., 2020). In general, for all samples, AAE_{MSOC} was found to be marginally higher than the respective AAE_{WSOC} . The largest variability in AAE_{MSOC} was observed among BFS smouldering (5.5-6.2) and flaming (5.3-6.0) samples, which was likely a result of the difference in combustion and heterogeneous vegetation distribution in the burnt BFS samples. The Svalbard peat (NOR) emission displayed AAE_{WSOC} in the range of 4.6-5.2, while its AAE_{MSOC} was constrained in the range of 4.9-5.2. In comparison to open BB, fresh RWC emissions having high soot contents (Mukherjee et al., 2024; wood stove 1 in Figure S7) exhibited much lower AAE for both WSOC and MSOC (2.2-2.4), while the wood combustions from the previous campaign (Ihalainen et al., 2019a; wood stove 2 in Fig. S7) displayed larger AAE_{WSOC} (5.8-6.8) and AAE_{MSOC} (5-5.6). The lower AAE values of RWC emissions indicates that BrC from high EC/OC emissions absorb light throughout the measured spectral range, as shown earlier by Saleh et al. (2018).

The MAC values obtained from UV-vis spectroscopy at 365 nm for fresh BB emitted MSOC (MAC_{365_MSOC}) were constrained between 0.46-1.48 $m^2 g^{-1}$ for the experiments. In comparison, MAC_{WSOC} was found to be in the range of 0.51-1.78 $m^2 g^{-1}$. These MAC_{365} values fall within the same range as observed for BBOA in previous literature (Park and Yu, 2016; Huo et al., 2018; Moschos et al., 2018; Cao et al., 2021). Among open BB experiments, the lowest MAC_{365_MSOC} were generally observed for fresh emissions from BFS



smouldering combustions, while flaming combustion of CP and SW were estimated to have the highest MAC_{365_msoc} (Table S4). The MAC_{365_msoc} for smouldering SG and SW emissions were in a similar range to the MAC_{OA_370} estimated by Vakkari et al. (2025). In general, flaming dominated BB emissions had higher MAC_{365_msoc} values compared to smouldering emissions (Fig. 5a). In addition, MAC values for high temperature wood log combustion in the modern European stove were significantly higher than for open BB. Fresh RWC emissions had MAC_{365_msoc} values in the range of $3.33\text{--}16.4\text{ m}^2\text{ g}^{-1}$, while the MAC_{365_wsoc} values ranged from $0.32\text{--}4.96\text{ m}^2\text{ g}^{-1}$. The correlation between the EC/OC ratios obtained for different experiments to the corresponding MAC_{msoc} also suggests that EC rich emissions contribute to stronger BrC light absorption (Fig. 5a). Formation of BrC internally mixed with soot particles in high temperature BB emissions has been shown to contribute to enhanced light absorption in past studies, due to lensing effect (Jacobson et al., 2001; Liu et al., 2015; Liu et al., 2017; Zhang et al., 2025), which supports our observation.

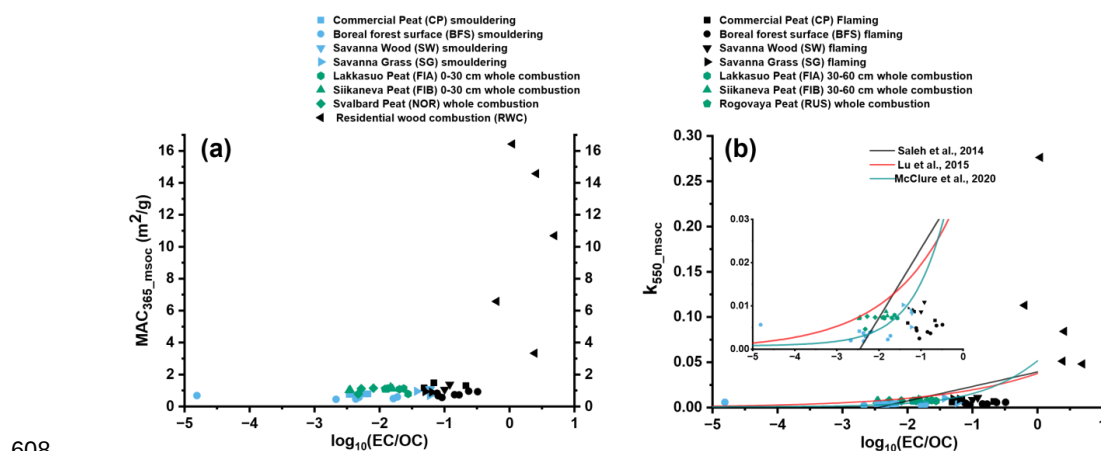


Figure 5: Relationships between MAC_{365_msoc} (a) and k_{550_msoc} (b) with EC/OC of open BB and RWC emissions.

We observe increasing MAC and k for high temperature combustions with higher EC/OC ratio

The imaginary refractive index measured at the middle of the visible light spectrum at 550 nm (k_{550}) for the BB MSOC varied across orders of magnitude and ranged between 0.002–0.011 (Fig. 5b, Table S4) in our experiment. Similar to MAC, the smouldering emissions exhibited lower k_{550_msoc} than the flaming dominated burns. As shown in previous studies (Saleh et al., 2018; Saleh et al., 2020) there seem to exist a continuum between the ratio of EC/OC in the fresh BB emission to the BrC light absorption, as we again observed



enhanced k_{550_msoc} for EC rich wood stove emissions (fig. 5b). Most of the fresh BB MSOC fell in the weak BrC regime in the “k-w” space (Fig. 6a), proposed by Saleh et al. (2018), while some of the wood stove combustion generated particles, in contrast, were in the strong BrC domain. These results fall in line with previous observations of high temperature (and high MCE) biomass combustions generating “darker BrC”, which exhibit much stronger light absorption compared to lower temperature open BB emissions (Saleh et al., 2018). The data points in k-w space obtained in this study matches reasonably well with previous literature, as shown in Figure 6a.

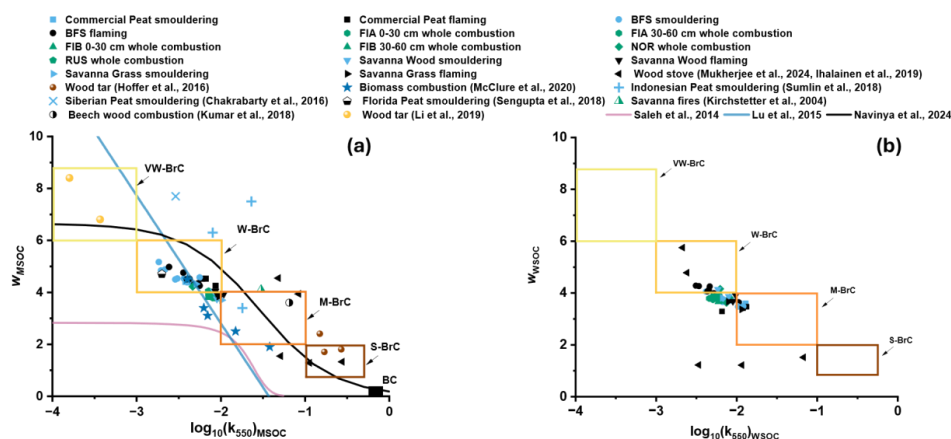


Figure 6: Relationship between w and k_{550} for (a) MSOC and (b) WSOC in fresh emissions from smouldering (light blue points), flaming (black points), full combustion (green points) of open BB performed in this study. Rectangles denote the different light absorbing BrC classes. Several data points from previous literatures (Hoffer et al., 2016; McClure et al., 2020; Sumlin et al., 2018; Chakrabarty et al., 2016; Sengupta et al., 2018; Kirchstetter et al., 2004; Kumar et al., 2018; Li et al., 2019; Saleh et al., 2014; Lu et al., 2015; Navinya et al., 2024) have been added in (a) for reference.

3.4 Particle size distributions and morphologies of chamber diluted primary emissions

Particle size distribution, morphology and effective density were measured for chamber diluted primary emissions. Smouldering combustion of BB generally emitted more particles on average when compared to flaming combustion (Fig. S8). Overall, there was little variability in the particle size distributions for flaming



635 combustion experiments, irrespective of the biomass type. The average particle geometric mean diameters
636 (GMD) ranged between 48 and 85 nm (Table 1). However, we observed clear variation in the emitted
637 particle size distribution among the smouldering combustions, likely due to the unique combustion behaviour
638 of individual samples with varying distributions of different types of biomass, such as surface vegetation,
639 litter and woody branches. Smouldering combustion of BFS gave rise to distinct bimodal size distributions
640 with the lower mode having a GMD of 30-40 nm, and the larger mode peaking around 100-175 nm. Some
641 replicates of smouldering emissions from CP and SW burning also exhibited bimodal size distribution.

642 The effective densities of primary particles inside the chamber varied between experimental replicates (Fig
643 S9), which can again be explained by the variability in the combustion and emissions between experiments.
644 The effective densities were largely independent of the particle diameter for all fuel type and combustion
645 conditions (Fig. S9), unlike soot-rich RWC emissions that exhibited size-dependent behaviour in our previous
646 work (Leskinen et al., 2014; Mukherjee et al., 2024). Due to the size-independent behaviour, we averaged over
647 all experimentally obtained density values and estimated average particle densities for each BB emissions
648 (Table S4). While CP and BFS combustion emitted particles had average densities of 1.1 g cm^{-3} and
649 $1.20(\pm 0.05) \text{ g cm}^{-3}$ for both smouldering and flaming dominated combustions, SW and SG emitted particles
650 were denser with average densities of 1.4 g cm^{-3} and $1.65(\pm 0.05) \text{ g cm}^{-3}$ for smouldering combustion
651 respectively. For flaming combustions, SW emitted primary particles had an average density of $1.45(\pm 0.05) \text{ g}$
652 cm^{-3} and SG particle density was $1.75(\pm 0.05) \text{ g cm}^{-3}$. Little to no size-dependency for particle density was
653 observed in our experiments suggesting near-spherical particle morphology with mobility exponents close to
654 3 (Leskinen et al., 2023; Corbin et al., 2023). SEM images of chamber diluted primary emissions (Fig. S10)
655 support these results, since mostly round organic particles were observed from the collected grids. These
656 findings agree with the measured chemical compositions that indicated that freshly emitted particles consisted
657 mostly of POC and very little EC, leading to formation of particles with a spherical morphology, so-called
658 tarballs. Tarballs have been previously observed in BB emissions (Chakrabarty et al., 2010; China et al., 2013;
659 Hoffer et al., 2016; Adachi et al., 2019) and atmospheric dilution and aging have been shown to aid the
660 formation of spherical viscous tarball BrC (Hennigan et al., 2011; Sedlacek et al., 2018). In chamber diluted
661 CP and BFS emissions, we observed mostly tarballs (Fig. S10a-d) but SG and SW emissions (Fig. S10e-h)

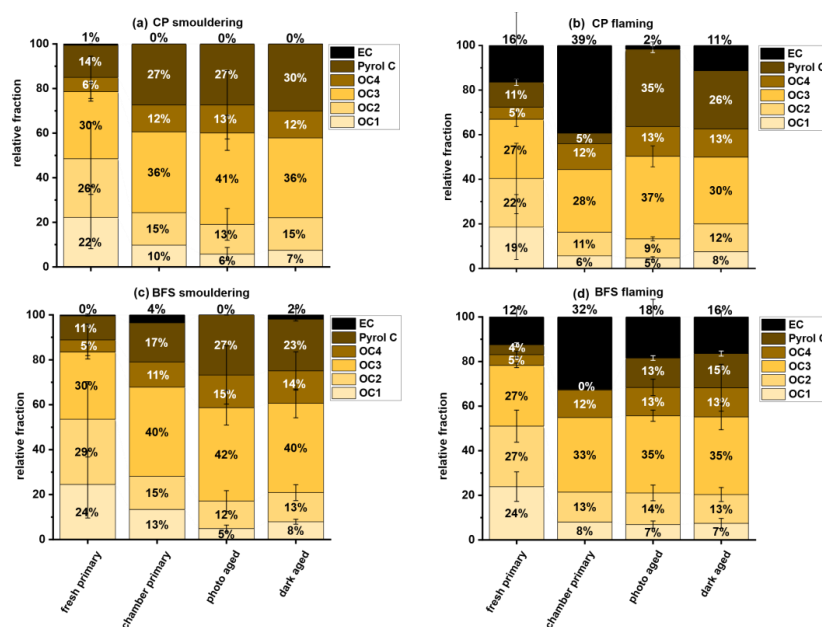


662 also had soot agglomerates that were partially coated or embedded with organics, as previously reported in
663 China et al. (2013). Notably, the morphology of fresh particles emitted from open BB was very different to
664 those emitted from a wood stove (Mukherjee et al., 2024), where the high temperature combustion formed
665 mostly chain-like fractal soot agglomerate structures which were mostly bare or with some organic inclusions
666 (Fig. S10i; China et al., 2013).

667

668 **3.5 Effect of chamber dilution on the properties of fresh emission aerosols**

669 The relative fractions of OC1, OC2, OC3, OC4, Pyrol C and EC in fresh emissions differed before and after
670 dilution in the chamber (Fig.7; Table S3), highlighting the effect of dilution on the overall chemical
671 composition of BB emissions. The relative fraction of the intermediate volatility (IVOC) and semi-volatile
672 organics (SVOC), namely OC1 and OC2 obtained from IMPROVE-A protocol (Ma et al., 2016) were lower
673 in the diluted chamber samples than in the fresh emissions, while the low volatile fractions from OC3 to PC
674 increased (Fig. 7). We might have over-estimated the OC1 fraction in fresh emission due to gas phase VOC
675 adsorption on the filter surface causing positive artifact (Kirchstetter et al., 2001). However, the significant
676 increase in low volatile OC fractions (OC3, OC4 and PC) in chamber samples indicate dilution related
677 evaporative loss of the volatile OC fractions in the chamber due to partitioning between gas and particle
678 phase (Calderon-Arrieta et al., 2024).



679

680 **Figure 7: The relative abundance of different OC fractions and EC measured in BB emissions from (a) fresh**
681 **emission, (b) fresh emission after dilution in chamber, (c) photochemically and (d) dark aged emission in the**
682 **chamber. Error bar indicates the variability among experimental replicates**

683 The partitioning of OC1 and OC2 fractions between gaseous and condensed phases is largely influenced by

684 the prevailing concentrations in the collected aerosol. Thus, a clear trend is seen between the share of low

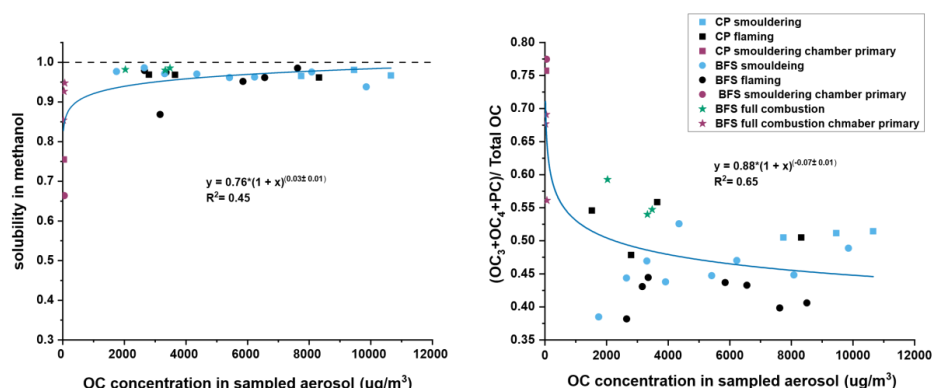
685 volatile OC fractions (OC3+OC4+PC) to the total OC and the collected aerosol OC concentration (Fig. 8).

686 The dilution and consequent evaporation of the most volatile OC also influences the solubility of the POC of

687 the fresh emission. Methanol extraction efficiency significantly decreased for the more diluted POC in the

688 chamber (Fig. S11b). This is in line with the change in volatility as the low volatile OC fraction, including

689 tarball BrC, is generally less soluble in solvents (Chakrabarty et al., 2023; Saleh et al., 2020).



690

691 **Figure 8: Effect of OC concentration on the partitioning of the OC fractions in BB emissions. Higher dilution in**
 692 **the chamber resulting in lower OC concentration result into larger fraction of low volatile OC fractions and**
 693 **lower solubility in methanol**

694

695 The increment in the low volatile OC fractions in chamber diluted primary emissions could also be seen from
 696 the abundance of spherical tarballs in the SEM images (Fig. S15). We postulate, based on indirect evidence
 697 and previous knowledge, that the fresh BB emission collected on the filters probably had poorly formed
 698 tarballs (consisting of higher OC1 and OC2 fractions) and higher methanol extractable OC. After the dilution
 699 related change in volatility distribution of the OC in the chamber (Calderon-Arrieta et al., 2024), methanol
 700 insoluble, highly viscous and spherical tarballs were formed in the chamber (Heninghan et al., 2011;
 701 Sedlacek et al., 2018; Adachi and Buseck, 2011). Interestingly, we observed a slight increase in the solubility
 702 of WSOC for chamber primary samples (Fig. S11a). This might be the effect of low OC amount filters
 703 collecting chamber diluted samples, leading to higher water to OC ratio during the extraction. Thus, the
 704 volume of water used for extracting chamber primary filter punches allowed the dissolution of sparingly
 705 water-soluble organics.

706

707 The change in particle volatility and solubility due to dilution in the chamber did not significantly affect the
 708 light absorption wavelength dependency of MSOC and WSOC. The AAE_{MSOC} and AAE_{WSOC} of chamber
 709 diluted samples ranged from 5.5-6.1 and 5-5.9, respectively. At the same time, the AAE_{MSOC} and AAE_{WSOC}



710 for the corresponding freshly emitted particles were 5.3-5.6 and 4.6-5.3, respectively. However, we observed
711 a 17% decrease in MAC_{365_MSOC} (Fig. S12) and a 25% decrease in MAC_{365_WSOC} (Fig. S13) in the chamber
712 primary samples compared to fresh emissions for BFS full combustion samples. This suggests that the
713 fraction of OC extracted by water and methanol in chamber diluted primary samples had smaller mass
714 absorption efficiency compared to those extracted from the fresh emissions. This can again be explained by
715 the formation of non-soluble dark BrC through the phenomenon termed as “darkening by volatilization” as
716 the low volatile fraction consists of stronger chromophores (Calderon-Arrieta et al., 2024).

717

718 **3.6. Effects of simulated atmospheric aging on properties of particulate organic** 719 **matter**

720 **3.6.1 Effects of aging on volatility and solubility of OC**

721 The relative OC fractions of chamber diluted primary particles and aged particles seemed to be similar (Fig.
722 7), suggesting that when it comes to the relative distribution of the OC volatility in BB emissions, the effect
723 of dilution in the environmental chamber outweighed the impact of photo and dark oxidative aging we
724 achieved in our experiments. After oxidative aging there were, however, further increments of the pyrolyzed
725 carbon (PC), which is the lowest volatility organic fraction. Oxidative aging is known to increase the
726 oxidation state and decrease the volatility of aged particles, making them more insoluble in organic solvents
727 (Saleh et al., 2020). Expectedly, we observed that both photochemical and dark oxidative aging decreased the
728 fraction of MSOC significantly (Figure S11b). On the other hand, the more oxygenated aged OC exhibited
729 slightly higher WSOC compared to POC in fresh emissions, and we observed a medium correlation ($R = 0.5$)
730 between water solubility and the atomic O:C ratio obtained from SP-AMS (Fig. S14b). The only exception to
731 this were the smouldering emissions of CP and BFS, which did not exhibit any significant change in water
732 solubility after dark or photochemical aging. This suggests that the chamber diluted POA emitted from
733 smouldering burns are more resistant to photochemical or dark oxidation, inadvertently hinting towards the
734 highly viscous tarball BrC. Tarballs are known to be resistant to chemical oxidation (Chakrabarty et al.,



2023) and we also observed the formation of more stable dark tarballs, especially in smouldering emissions, due to oxidative aging in the chamber (Fig. S15).

3.6.2 Effects of aging on organic matter density

Oxidative aging, both in photochemical and high NO_x dark conditions, led to a marginal increment in particle densities for BFS emissions. For BFS smouldering emission, the effective density after photochemical aging increased to $1.30 \pm 0.05 \text{ g cm}^{-3}$ while no significant change was observed between primary and dark aged particle densities. For flaming dominated combustion, the particle densities increased under both photochemical ($1.30 \pm 0.09 \text{ g cm}^{-3}$) and dark aging conditions ($1.30 \pm 0.07 \text{ g cm}^{-3}$). For CP, the effective density of high NO_x dark aged particulate emissions remained unaltered at 1.1 g cm^{-3} , while photochemically aged emissions exhibited higher density. For SW emissions, both photochemical aging and low NO_x dark aging appeared to have a negligible impact on overall particle effective densities (Table S4). In contrast, particle effective densities for flaming and smouldering combustion of SG emissions seemed to be lower after undergoing photochemical and dark aging (Table S4).

3.6.3 Effects of aging on organic matter chemical composition

Due to low OC loading on the filters collected from the chamber, we were only able to assign roughly 120-600 unique elemental formulae of the most abundant chemical compounds from FT-ICR MS analyses of the photo and dark aged samples of CP and BFS, while signals arising from other compounds were below the detection threshold ($S/N \geq 6$) (Fig. S16, S17). Therefore, we were unable to perform a one-to-one comparison of the chemical compositions of the fresh emission with the chamber-aged samples. However, we could conclude that CH and CHO were the most abundant compound classes in the chamber diluted primary and aged samples for both CP and BFS. We obtained elemental ratios (O:C, H:C and N:C) of the chamber diluted primary and aged samples from the SP-AMS measurement to characterize the variability in the bulk composition of the OC. The resulting Van-Krevelen plots (Fig. S18) showed that the chamber diluted primary emission had H:C values close to 2.0, while the O:C and N:C values ranged between 0.10-0.25 and 0.013-0.027, respectively. We observed an increment in the O:C ratios in the chamber aged emission while



761 the H:C ratio decreased, as in some previous studies (Lambe et al., 2011). On the other hand, after high NO_x
762 dark aging, the overall N:C ratio did not display a similar increase, probably due to the fast degradation of
763 the nitroaromatics formed inside the chamber at the time of filter collection at the end of the experiments.

764 3.6.4 Effects of aging on optical properties

765 Photochemical aging increased AAE and the wavelength dependency of $k(w)$ of the BB emissions for both
766 MSOC and WSOC. The light absorption shifted towards smaller wavelengths (fig. S19) while the k at 550
767 nm either remained unchanged or decreased (Fig. 9). w_{WSOC} and w_{MSOC} varied between 3.3-3.7 and 4.3-4.5
768 respectively for fresh emission of CP, while it increased to 4.0-5.1 and 4.0-6.4 after photochemical aging.
769 The largest increase in w_{MSOC} was observed for flaming combustion of SG, where w_{MSOC} increased from 3.9
770 to 6.3, while k_{550_MSOC} decreased an order of magnitude. All the photochemically aged WSOC and MSOC
771 samples in this study lied in the w-BrC to vw-BrC region of the k-w space. This behaviour is in line with the
772 effects of photobleaching and photolysis during OH exposure in the presence of UV lights. Increasing
773 photooxidation has been shown to cause fragmentation reactions after functionalization (Kroll et al., 2015;
774 Saleh et al., 2020). OH• has been shown to cleave polyaromatic chains and give rise to short unsaturated
775 chemical moieties that absorb light at the UV range of the spectrum. On the other hand, photochemical
776 aging of RWC emissions in oxidation flow reactor either decreased (for samples from Ihalainen et al., 2019a)
777 or increased (Mukherjee et al., 2024) the k_{550_msoc} , suggesting that chemical composition of the precursor
778 organic molecules and the reaction pathways are consequential to secondary BrC optical properties (Lambe
779 et al., 2013; Sumlin et al., 2017; Hems et al., 2020).

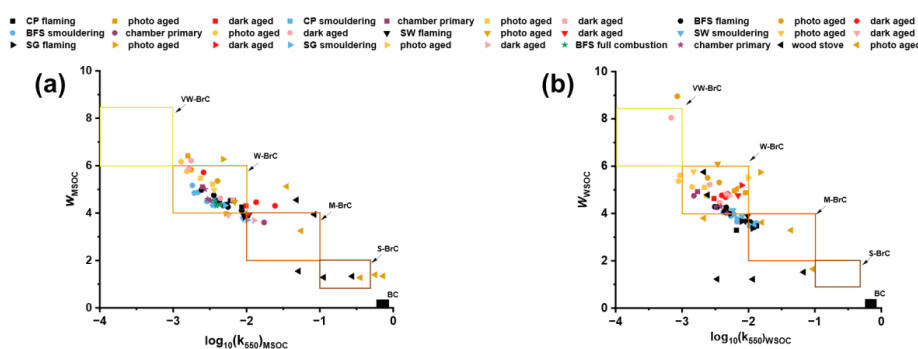
780

781 The effect of dark aging on the optical properties was more non-trivial in our study. We noted that for
782 flaming dominated emissions, the NO_x dominated dark aging led to higher k_{550} values and decreasing w
783 values (fig. 9) in accordance with Saleh (2020). This means an increase in light absorption towards the
784 visible wavelength range, which was also seen in the absorption spectra (fig. S19). On the other hand, dark
785 aging seemed to have very minor effect on the smouldering dominated emissions. One explanation for this
786 observation could be the prevalence of tarballs in the smouldering dominated emissions of both peat and
787 other biomass samples, which have been shown to be quite resistant to oxidative aging (Chakrabarty et al.,



2023). Additionally, nitroaromatics are reactive, and certain fractions of them may have degraded at the end of the chamber experiments or during filter collection from the chamber, making it difficult to distinguish the effects on optical properties.

791



792

793 **Figure 9: k-w space for all experimental data included in our study. Black, blue and green colours denote fresh**
794 **emissions from flaming, smouldering and full combustion respectively. Purple dots denote primary emission in**
795 **chamber without any aging (chamber primary). Dark yellow and red points denote photochemical and dark**
796 **aging of flaming dominated emissions respectively, while light yellow and red dots represent respective**
797 **photochemical and dark aging of smouldering dominated emissions.**

798 WSOC light absorption in “k-w space” (fig. 9b) suggests that the water soluble BrC in fresh and aged
799 emissions of RWC had lower k_{550} values and therefore less strongly absorbing chromophores than the
800 corresponding MSOC samples. Certain dark aged samples, especially those for flaming BFS and SG
801 emissions, also seemed to contain stronger chromophores in MSOC than in WSOC. These observations
802 suggest that less polar organics, which are soluble in methanol but not in water, contribute to higher light
803 absorption in these BB emissions. In addition, it should be noted that MSOC light absorption in the aged
804 emissions might be underestimated, because fractions of the aged organics were non-soluble (Fig. S12b) and
805 might have contained even stronger chromophores. In future, online measurements of OA absorption at
806 different wavelengths is needed to confirm the observed effects of photochemical and dark aging.

807

808 4. Conclusions:



809 In this work, we estimated OC and EC emission factors and organic aerosol optical and chemical properties
810 for different Eurasian open BB emissions, separately for flaming and smouldering dominated combustion.
811 These characteristics were compared with two other major source-specific BB emissions; residential wood
812 combustion and African savannah biomass burning. Open BB particulate emissions are generally dominated
813 by OC and specifically smouldering conditions emit higher OC than flaming combustion, as expected. Open
814 BB OC emission factors ranged from 1.39 to 89.9 g kg⁻¹ fuel.

815 Fraction of WSOC varied in the emissions from different biomasses, whereas nearly all of the
816 emitted fresh OA was methanol soluble (>92%). Therefore, we used the optical properties of MSOC as a
817 proxy for bulk BrC optical properties. The imaginary refractive index at 550 nm for MSOC (k_{550_MSOC}) of the
818 fresh organic aerosols varied between 0.002 and 0.011, and fell into the category of weakly absorbing BrC,
819 as classified by Saleh (2020). The MAC_{365_MSOC} ranged between 0.46 and 1.48 m² g⁻¹ and was observed to
820 follow a continuum with respect to aerosol EC/OC-ratio. Consequently, flaming emissions exhibited higher
821 MAC_{MSOC} (can be also called MAC_{OA}) than smouldering emissions. In addition, the biomass-type influenced
822 light-absorption of OA, with peat emissions showing the highest MAC_{OA} , most likely due to their high
823 abundance of nitroaromatic compounds. AAE_{MSOC} in the wavelength of 300-550 nm ranged between 4.2-6.2,
824 indicating that the light absorption was strongly dominated by BrC chromophores that absorb at shorter
825 wavelengths.

826 Further characterization of the BB emissions was carried out using an environmental chamber
827 to simulate the effects of atmospheric dilution and oxidative aging on aerosol properties. Irrespective of
828 biomass type, smouldering combustion generated higher particle number concentrations than flaming
829 dominated burns. Tarball morphologies of variable sizes were observed for all open BB emissions and in all
830 studied conditions. OC1 and OC2 fractions, that are the most volatile OC fractions and overlap with SVOC
831 and IVOC, decreased in the chamber compared to fresh emissions, due to particle-to-gas partitioning in the
832 chamber. This dilution induced change of OA volatility distribution resulted in higher fractions of OC3, OC4
833 and PC (SVOC, LVOC and ELVOC) and potentially formed stronger BrC chromophores (Calderon-Arrieta
834 et al., 2024) in the chamber primary samples. However, we observed a slight decrease in MAC_{MSOC} for



835 wavelengths higher than 365 nm. This contrasting finding is likely because our solvent extraction method
836 cannot resolve the lowest volatility fraction of OA in the chamber samples.

837 After photochemical and dark oxidation, the O:C ratio of the particles increased and resulted
838 in increasing ELVOC fractions (PC in particular), making it more insoluble in methanol. Dark aged CP and
839 BFS smouldering emissions in particular had the lowest methanol solubility, probably due to the high
840 abundance of tarballs with high viscosity OA (Chakrabarty et al., 2023). On the other hand, WSOC fraction
841 moderately increased after aging, likely due to the observed higher oxygenation and consequent polarization
842 of the compounds. For the dark-aged samples, absorption increased for all wavelengths, suggesting the
843 formation of stronger light-absorbing oxygenated compounds and/or nitroaromatics. For photochemically
844 aged samples, MAC₃₆₅ increased but visible wavelength absorption, i.e. k_{550} decreased, probably due to
845 photobleaching and breaking down of BrC chromophores. In future, the OA offline optical characterization
846 should be accompanied by online aerosol optical measurements, to estimate light-absorption of non-soluble
847 OA.

848 Overall, our results highlight that boreal forest surface fires as well as boreal and arctic peat fires are
849 potentially important sources of BrC into the atmosphere. Further, we show that atmospheric dilution and
850 oxidative aging influence BrC properties. This dynamic property of BrC should be considered in global
851 climate models to correctly estimate direct radiative forcing of biomass burning emissions.

852

853 **Data availability:**

854 Essential data used in this study are available at FAIR-aligned data repository Zenodo via
855 <https://doi.org/10.5281/zenodo.15647186> with Creative Commons Attribution 4.0 International license
856 (Mukherjee et al., 2025). Detailed mass spectrometry data are available upon request from the corresponding
857 authors.

858

859 **Author contribution:**



VV, AnV, OS, AH, PYP, MI, KK and AB contributed in designing the study as well as supervising the whole experimental campaign. KK and MR arranged for the arctic-boreal biomasses, whereas KJ, SJS, PGVZ and VV collected, sorted and transported savanna biomasses to the study site in Finland. Measurements were performed by AM, MaS, AH, MI, PYP, LMFB, DL, TuK, LV, AB, MuS and VV. Offline filter analyses, including solvent extraction and thermal–optical analyses were performed by AM and VL, with supervisions from NK. SEM images were captured by HK. FT- ICR MS analyses of filter samples and subsequent data analyses and plotting were done by CPR, TiK, VHN and JJ. Other data analyses were performed by AM, MaS, TuK, AH, LB and LV. The manuscript was prepared by AM with help from AH, OS, MR, KK, MaS, LB, VN, TK, CR, JJ, HT and AkV. Fundings for this work were acquired by OS, AkV, AnV, VV and HT.

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