



1 **Drivers and CO₂ flux budgets in a Sahelian *Faidherbia albida* agro-silvo-pastoral parkland:**
2 **Insights from continuous high-frequency soil chamber measurements and Eddy**
3 **Covariance.**

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29 **Highlights:**

- 30 • Long-term high frequency CO₂ flux measurements using automated static
31 chambers in a Sahelian *F. albida* parkland.
- 32 • Empirical gap-filling and flux partitioning methods validated against Eddy
33 Covariance GPP.
- 34 • Fluxes peaked during the rainy season in both FS and Sh, driven mainly by soil
35 moisture and leaf area.
- 36 • *F. albida* trees enhance CO₂ fluxes under canopies ("fertile island" effect) and
37 account for ~50% of annual ecosystem GPP.



38 **ABSTRACT:**

39 Agroforestry systems — combining trees with crops and/or livestock — are increasingly
40 promoted as sustainable and climate-resilient land-use strategies. Despite their widespread
41 presence in the Sahel, experimental data on their potential as carbon sinks are scarce. This study
42 presents a full-year, high-frequency dataset of CO₂ fluxes in a Sahelian agro-silvo-pastoral
43 parkland dominated by *F. albida*, located in Senegal's groundnut basin. CO₂ fluxes were
44 continuously measured using automated static chambers, allowing the quantification of soil and
45 crop respiration (R_{ch}), gross primary production (GPP_{ch}), and net carbon exchange (FCO_{2ch})
46 under both full sun and shaded (under tree canopies) environments.

47 Seasonal patterns of CO₂ fluxes were similar in both environments, with peaks during the rainy
48 season. R_{ch} and GPP_{ch} were significantly higher under tree canopies, indicating a 'fertile island'
49 effect. CO₂ flux variability was primarily driven by soil moisture and leaf area index. Chamber-
50 based GPP estimates closely matched those from Eddy Covariance measurements. On an annual
51 scale, *F. albida* trees contributed approximately 50% of total ecosystem GPP, with a carbon use
52 efficiency of 0.48. Net annual CO₂ exchange was estimated at -1.4 ± 0.02 and -1.8 ± 0.01 Mg C-
53 CO₂ ha⁻¹ using chamber and Eddy Covariance methods, respectively. These findings underscore
54 the role of *F. albida*-based agroforestry systems as effective carbon sinks in Sahelian landscapes,
55 supporting their potential contribution to climate change mitigation.

56 **Keywords:** Sahelian agro-silvo-pastoral systems, CO₂ fluxes, automated static chambers, Eddy
57 Covariance, 'fertile island effect' of trees, carbon budgets.



58 1. Introduction

59 Plant photosynthesis and respiration —both autotrophic (plant) and heterotrophic (microbial)—
60 are fundamental processes driving carbon dioxide (CO₂) fluxes in terrestrial ecosystems
61 (Lambers et al., 2008; Raich et al., 2014; Reichle, 2020). Accurate quantification of these processes
62 is critical for assessing ecosystem carbon (C) sink potential (Baldocchi, 2020), particularly for
63 informing climate-smart land management strategies.

64 To capture these processes at the ecosystem scale, the Eddy Covariance (EC) technique has
65 emerged as a transformative method, enabling continuous and high-frequency CO₂ flux
66 measurements (Baldocchi, 2003, 2008). Extensive EC networks in Europe (Stojanović et al.,
67 2024), Asia (Yu et al., 2011), and the Americas (Chu et al., 2021) have significantly advanced our
68 understanding of the global C cycle. In contrast, sub-Saharan Africa remains critically
69 underrepresented (Bombelli et al., 2009; Houghton & Hackler, 2006; Williams et al., 2007).
70 Although some studies have used EC (Ardö et al., 2008; Brümmer et al., 2008; Merbold et al., 2009;
71 Tagesson et al., 2016), static chambers (Assouma et al., 2017; Owusu et al., 2024; Rosenstock et
72 al., 2016; Wachiye et al., 2020), or modeling approaches (Agbohessou et al., 2023, 2024; Delon et
73 al., 2019; Rahimi et al., 2021), they remain sparse and methodologically heterogeneous, limiting
74 comparability and regional C budget integration.

75 Among these underrepresented landscapes, agroforestry systems in the Sahel— particularly
76 agro-silvo-pastoral systems (ASPS) that combine trees, crops, and livestock— are increasingly
77 promoted for sustainable land management and climate resilience (Cardinael et al., 2021; Gupta
78 et al., 2023; Mbow et al., 2014; Stetter & Sauer, 2024). However, the structural and functional
79 heterogeneity of these systems poses significant challenges for accurately quantifying and
80 upscaling C fluxes. *Faidherbia albida*, a keystone agroforestry tree species in these ASPS (Leroux
81 et al., 2022; Lu et al., 2022), is of particular interest due to its reverse phenology, capacity to
82 enhance soil fertility and crop yields (Bayala et al., 2020; Roupsard et al., 2020; Sileshi et al., 2016;
83 2020). Yet, its functional role in modulating both the magnitude and seasonal dynamics of CO₂
84 fluxes remains poorly understood.

85 Addressing this knowledge gap requires integrated approaches capable of capturing both
86 aggregate and component-specific CO₂ fluxes. While EC remains the gold standard method for CO₂
87 flux measurements at the landscape scale (Baldocchi, 2003), it captures net ecosystem exchange
88 (NEE) as an aggregate signal, without separating the contributions from individual compartments
89 such as soil, crops, and trees. This limits its utility for disentangling processes and attributing
90 sources in heterogeneous systems like ASPS. Automatic static chambers provide a valuable
91 complement to EC, as they enable continuous, high-frequency measurements at finer scales and
92 at the level of specific ecosystem components. This approach facilitates component-specific
93 quantification of CO₂ fluxes, particularly from soil and crop compartments (Luo & Zhou, 2006;



94 Denmead, 2008; Zaman et al., 2021). When combined with EC, this dual-method approach
95 strengthens source attribution and improves the upscaling of fluxes across complex agroforestry
96 landscapes.

97 This study presents one of the first integrated quantification of CO₂ fluxes in a Sahelian ASPS
98 dominated by *F. albida*, combining EC and automatic static chambers.

99 Specifically, we aim to (1) conduct year-round, high-frequency *in situ* CO₂ flux measurements
100 from soil and crops using automated static chambers; (2) partition the net CO₂ fluxes (FCO₂ch)
101 into respiration (R_{ch}) and photosynthesis (GPP_{ch}); (3) investigate the environmental drivers of
102 fluxes and the spatial variability linked to tree presence; and (4) compare chamber-based flux
103 estimates with ecosystem-scale measurements derived from the EC method.



104 2. Materials and methods

105 2.1. Site description

106 The study was conducted in the agroforestry parkland of Sob village (Niakhar municipality, Fatick
107 region), located in the groundnut basin of Senegal, within the Sahelo-Sudanian climatic zone of
108 West Africa (Fig. 1). The climate is characterized by a long dry season (8–9 months) with high
109 temperatures and strong diurnal variations, and a short rainy season from late June to mid-
110 October (Delaunay et al., 2018).

111 Soils are locally known as “Dior” and classified as Arenosols (IUSS Working Group WRB, 2022).
112 The topsoil has low organic matter (<1%) and phosphorus (<3 mg kg⁻¹), a sandy texture (>85%
113 sand), and an acidic pH (Malou et al., 2021; Siegwart et al., 2022). Rainfed agriculture
114 predominates. The main cropping system includes pearl millet (*Pennisetum glaucum* L.) and
115 groundnut (*Arachis hypogaea* L.) in biennial rotation, with occasional intercropping of cowpea
116 (*Vigna unguiculata* L.).

117 The site hosts the 'Faidherbia Flux' station (14°29'44.916"N; 16°27'12.851"W; FLUXNET ID: SN-
118 Nkr), a long-term research platform for monitoring ecosystem services in agroforestry systems.
119 It is dominated by *F. albida*, a nitrogen-fixing, reverse-phenology tree with deep roots accessing
120 groundwater (Roupsard et al., 1999). The tree density is ~13 trees ha⁻¹, with canopies covering
121 ~10% of the soil surface (Roupsard et al., 2020). The EC tower is installed at 20 m height,
122 approximately 12.5 m above the canopy. The study field is a typical 'bush field', characterized by
123 low soil fertility, no mineral fertilization, and off-site export of crop residues and manure (Malou
124 et al., 2021).

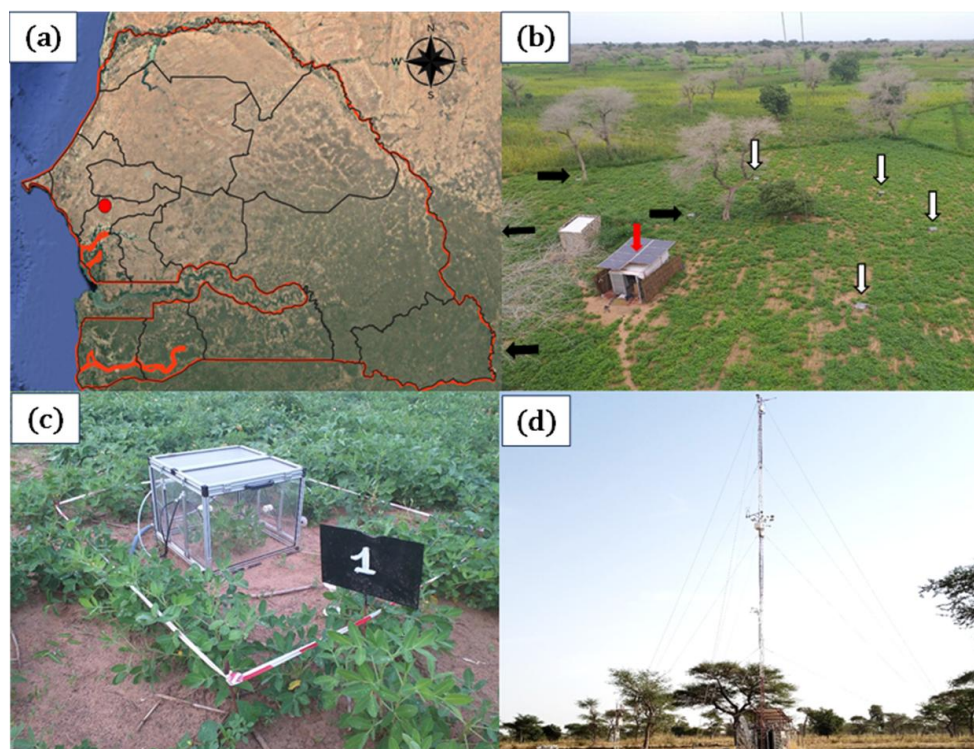


Fig. 1: Study area.

(a) geographical location of Sob, Groundnut basin, Senegal (Map data © Google Earth, 2025), (b) overview (image from the Eddy Covariance tower located in the same bush-field) of the *Faidherbia albida* parkland during the rainy season, depicting groundnut crops with bare soil in the inter-row, *F. albida* trees (defoliated during the rainy season, average height = 13m) and location of the chambers under the Shade of trees (horizontal black arrows; N=4) and in Full sun (vertical white arrows; N=4); The shelter (red arrow) with solar panels is to fit the analyser, automation and batteries (c) automatic chamber enclosing a groundnut plant (during the rainy season) or bare soil (during the dry season), (d) Eddy Covariance (EC) tower (measurement height = 20 m) during the dry season.



134 *2.2. Experimental setup*

135 *2.2.1. CO₂ flux measurements in automatic chambers*

136 Continuous net CO₂ fluxes (FCO₂ch) from soil and groundnut plants were measured over a full
137 phenological year (June 17, 2021 – June 17, 2022) using eight automated static chambers
138 (50×50×50 cm), each enclosing one groundnut plant. Four chambers were installed in full sun
139 (FS), at least 20 m from trees, and four under *F. albida* canopy shade (Sh). The chambers were
140 transparent, custom-built (Duthoit et al., 2020), and installed on metal bases embedded 10 cm
141 into the soil one month prior to measurements.

142 During the rainy season (June–November), groundnut coexisted briefly with spontaneous weeds
143 until weeding (mid-July), after which chambers contained only groundnut. Post-harvest (early
144 November), chambers remained bare while surrounding plots experienced weed regrowth.

145 CO₂ concentrations were measured at 1 Hz using a Picarro G2508 gas analyser (Picarro Inc., Santa
146 Clara, CA, USA) (Fleck et al., 2013; Reum et al., 2019; Valujeva et al., 2022). A fully automated
147 system was built for sequential half-hour flux measurements (alternating FS and Sh).
148 Measurement duration was 15 min per chamber in the dry season, reduced to 5 min during the
149 rainy season to limit condensation effects.

150 *2.2.2. CO₂ flux measurements by Eddy Covariance*

151 The EC system (Li-COR SMARTFLUX®, including a Gill MasterPro 3D sonic anemometer and a LI-
152 7500 RS open path CO₂ and H₂O gas analyser) was mounted at a height of 20 m on a 30m mast,
153 above *F. albida*. It continuously monitored net CO₂ exchange from the ecosystem. Raw data were
154 collected at 20 Hz frequency and post-processed from binary files using the advanced mode of the
155 EddyPro® v7.0, with standard corrections and procedures: sonic tilt correction (double rotation),
156 block averaging, covariance maximisation for time lag, and WPL correction (Webb et al., 1980).
157 Quality control followed Foken et al. (2004) and Vickers & Mahrt (1997); random uncertainty was
158 estimated per Finkelstein & Sims (2001). Spectral corrections were applied according to
159 Moncrieff et al. (1997, 2004). Footprints were computed according to Kormann and Meixner
160 (2001), using the FREddyPro R package (Xenakis, 2016), indicated a ~1 ha source area covering
161 the entire field. Gap-filling and flux partitioning were conducted using ReddyProc (Wutzler et al.,
162 2018), applying the daytime partitioning approach of Lasslop et al. (2010).

163 *2.2.3. Ancillary measurements*

164 Environmental and vegetation variables were monitored continuously throughout the study.
165 Global radiation (R_g) was estimated from photosynthetically active radiation (PAR) using a Skye
166 sensor (averaged over 30-min intervals). NDVI of crops under full sun was recorded semi-hourly



167 by a calibrated downward-facing sensor installed at 20 m height (Pontauiller et al., 2003),
168 processed following Soudani et al. (2012), and used to estimate LAI time series for groundnut,
169 weeds, and cowpea based on end-of-season field LAI measurements in six 15 m² plots (as in
170 Rounsard et al., 2020).

171 Rainfall was recorded by an automatic weather station (CR1000 with TE525MM rain gauge,
172 Campbell Scientific), and soil volumetric water content (VWC) and temperature (T_{soil}, at 6 cm
173 depth) were monitored using TOMST® TMS-4 sensors, benchmarked prior to field deployment
174 inside and outside the chambers (Wild et al., 2019). Air temperature (T_{air}) was recorded inside
175 each chamber at 15 cm above ground, all at 5-min intervals. These measurements contribute to
176 the SoilTemp global database (Lembrechts et al., 2020, 2022).

177 Groundnut development was tracked weekly by counting leaves in each chamber. Total
178 groundnut LAI (LAI_{ch}) was then derived from average single-leaf area and chamber surface.

179 A detailed description of the data used in this study is provided in Supplement S1 (Table S1.1).

180 2.3. Data processing

181 2.3.1. Flux calculation

182 Net CO₂ fluxes (FCO_{2ch}, in μmol CO₂ m⁻² s⁻¹) from the chambers were calculated from the linear
183 change in CO₂ concentration over time (ΔC/Δt) using the Eq.1.

$$184 \text{ FCO}_{2\text{ch}} = \left(\frac{P}{R T_k}\right) \left(\frac{V}{A}\right) \left(\frac{\Delta C}{\Delta t}\right) \quad (\text{Eq. 1})$$

185 where P is atmospheric pressure (101 325 N m⁻²), R is the ideal gas constant (8.31 N m mol⁻¹ K⁻¹),
186 T_k is air temperature inside the chamber in Kelvin, V (0.125 m³) is the total system volume
187 (chamber, tubing, analyser cavity, pump, and water trap), and A (0.25 m²) is the chamber
188 footprint. The slope ΔC/Δt was obtained via linear regression (Duthoit et al., 2020).

189 Mean FCO_{2ch} values were computed separately for the four replicate chambers in full sun (FS)
190 and under *F. albida* shade (Sh). By convention, negative values indicate net CO₂ uptake
191 (photosynthesis), and positive values indicate net CO₂ release (respiration).

192 2.3.2. Quality control of chamber-based CO₂ flux measurements

193 The quality of chamber-based CO₂ flux measurements was assessed using the coefficient of
194 determination (R² ≥ 0.8) of the linear increase in CO₂ concentration during chamber closure. The
195 minimum detectable flux (MDF) was then calculated following Nickerson (2016) (Eq.2). The MDF
196 defines the flux detection threshold, below which data are considered unreliable due to
197 instrument sensitivity and sampling constraints (Zaman et al., 2021). In this study, the MDF was
198 ±0.0004 μmol CO₂ m⁻² s⁻¹.

199



200
$$\text{MDF} = \left(\frac{A_a}{t_c(\sqrt{t_c/p_s})} \right) \left(\frac{VP}{ART} \right) \quad (\text{Eq. 2})$$

201 where A_a is the analytical precision of the Picarro analyser (0.6 ppm; Picarro Inc., 2015), t_c the
202 closure time (s), p_s the sampling frequency (1 Hz), V the chamber volume, P the atmospheric
203 pressure (101 325 N m⁻²), A the chamber footprint, R the gas constant (8.3 N m mol⁻¹·K⁻¹), and T
204 the air temperature in Kelvin.

205 Following this quality control, fluxes were partitioned (Section 2.3.3) and gap-filled (Section
206 2.3.4).

207 2.3.3. Partitioning of chamber-based CO₂ fluxes

208 The net CO₂ fluxes (FCO₂ch), averaged from four chambers per environment (FS and Sh), were
209 partitioned into two components according to Eq. 3 (Reichstein et al., 2005).

210
$$\text{FCO}_2\text{ch} = \text{Rch} + \text{GPPch} \quad (\text{Eq. 3})$$

211 Rch includes heterotrophic respiration (Rh) from soil and other autotrophic respiration (Ra) from
212 groundnut plants and roots of *F. albida* (Ra Groundnut + Ra tree below-ground). Rch is always
213 positive (Rch > 0). GPPch (Gross Primary Productivity) represents the photosynthetic CO₂ uptake
214 by the groundnut plants and is negative during the day (GPPch < 0), and zero at night, when
215 FCO₂ch = Rch.

216 Half-hourly FCO₂ch fluxes were partitioned as follows: (1) an Arrhenius-type function (Lloyd &
217 Taylor, 1994) was fitted between nocturnal Rch and T_{soil} during nighttime periods, for each 5-days
218 throughout the time series (Eq. 4). This empirical formulation is based on several key
219 assumptions. First, the relationship between nocturnal respiration and soil temperature is
220 assumed to follow an exponential response, reflecting the temperature sensitivity of respiration
221 processes. Second, the model assumes temporal stability of the respiration–temperature
222 relationship between night and day, allowing diurnal respiration to be extrapolated from fitted
223 parameters in Eq.4 and daytime T_{soil}. Third, we assumed that no abrupt changes in substrate
224 availability or soil moisture occur between day and night — conditions that could otherwise
225 disrupt the temperature–respiration relationship. Third, it is assumed that no abrupt changes in
226 substrate availability or soil moisture occur between night and day — conditions that could
227 otherwise decouple respiration rates from temperature. These assumptions are widely applied in
228 CO₂ flux partitioning approaches (Reichstein et al., 2005; Lasslop et al., 2010). (2) Diurnal Rch
229 was estimated by applying the Lloyd & Taylor function, previously calibrated on nocturnal data,
230 to the corresponding daytime T_{soil} measurements for each 5-day interval. (3) GPPch was
231 subsequently derived as the residual component of the net CO₂ flux during the day, according to:

232

233



234 **nocturnal Rch** = $R_{\text{ref}} \cdot \exp \left[E_0 \left(\frac{1}{T_{\text{ref}} - T_0} - \frac{1}{T_{\text{soil}} - T_0} \right) \right]$ (Eq. 4)

235 where R_{ref} ($\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$) is a fitted parameter representing the base respiration at the
236 reference temperature [T_{ref} (K), (set at 288.15 K)]. E_0 (K) is the temperature sensitivity (set at
237 250 K), T_{soil} (K) the soil temperature (K), and T_0 (K) is kept constant at 231.13 K, according to
238 Lloyd & Taylor (1994).

239 **GPPch** = **diurnal FCO₂ch** – **diurnal Rch** (Eq. 5)

240 where diurnal FCO₂ch and diurnal Rch represent the daytime net CO₂ fluxes and respiration in
241 $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$, respectively.

242 2.3.4. Gap-filling procedure

243 Missing Rch data were gap-filled using the model derived from Eq. 4 (Lloyd & Taylor, 1994). Prior
244 to gap-filling GPPch, raw data were standardised by LAI to reduce variability between chambers
245 due to differences in leaf surface area (Eq. 6). A light-response model was then fitted to the
246 standardised GPPch data, every 5-day period, to gap-fill missing values. The model is based on a
247 rectangular hyperbolic function that describes the relationship between photosynthetic CO₂
248 uptake and incoming global radiation (Rg) (Eq. 7). It corresponds to a Michaelis–Menten-type
249 light-response curve, commonly used in ecosystem carbon exchange studies (Falge et al., 2001;
250 Lasslop et al., 2010).

251 **GPPch.stand** = $\frac{\text{GPPch}}{\text{LAIch}} * \text{LAI.field}$ (Eq. 6)

252 where GPPch.stand ($\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$) is the standardised GPPch. LAIch and LAI.field ($\text{m}^2 \text{ leaves}$
253 $\text{m}^{-2} \text{ soil}$) represent the groundnut LAI inside the chambers and the groundnut + weeds + cowpea
254 LAI for the whole field, respectively.

255 **GPP** = $\frac{\alpha \beta R_g}{\alpha R_g + \beta}$ (Eq. 7)

256 where α ($\mu\text{mol CO}_2 \text{ J}^{-1}$) represents the light use efficiency of the groundnut plants inside the
257 chambers, and refers to the initial slope of the light-response curve, β ($\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$) is the
258 maximum CO₂ uptake rate by the groundnut plants at light saturation, and Rg the global radiation
259 (W m^{-2}).

260 2.3.5. Comparing chamber-based (Ch) and Eddy Covariance (EC) methods

261 Chamber measurements were upscaled to field-level CO₂ fluxes and compared with EC-derived
262 fluxes. Before comparison, a correction was applied (Eq. 6) to account for differences in LAI
263 between chambers (LAIch) and the field (LAI.field), due to the presence of cowpea and weeds in
264 the field but not in the weeded chambers.

265 Upscaling considered tree cover, with FS and Sh chamber fluxes weighted at 90% and 10%,
266 respectively. Rch.stand and GPPch.stand, representing chamber-based respiration and



267 photosynthesis at field scale. These fluxes were compared, on a half-hourly basis, to EC-derived
268 Reco.EC and GPP.EC (S3, Table S3.1). The November–December transition period was excluded
269 due to weed-driven uncertainties after groundnut harvest.

270 During the rainy season (*F. albida* leafless), GPP.EC represented ground vegetation (groundnut,
271 cowpea, weeds), while Reco.EC included autotrophic respiration from all vegetation (including
272 trees), and heterotrophic respiration (Reco.EC = R_a tree below-ground + R_a tree above-ground
273 + R_a groundnut + R_a cowpea + R_a weeds + R_h). R_{ch} .stand could not be fully upscaled to the field
274 due to uncertainty in its partitioning between R_a and R_h . R_{ch} .stand accounted only for R_a tree
275 below-ground, R_a groundnut, and R_h .

276 In the dry season (leafy trees, bare soil), GPP.EC reflected tree photosynthesis only (GPP tree),
277 while GPPch.stand was nil. Reco.EC included R_a tree (above- and below-ground) and R_h .
278 R_{ch} .stand, measured on bare soil represented only R_a tree below-ground + R_h .

279 2.3.6. Contribution of trees to full ecosystem respiration and photosynthesis

280 During the dry season, when the trees (*F. albida*) maintained their foliage, a comparison between
281 chamber and EC measurements allowed for the estimation of the contribution of the above-
282 ground tree compartments to total ecosystem respiration (S3, Table S3.1). Based on this estimate,
283 total tree respiration (R_a tree) was then calculated under the assumption that the tree root
284 systems (R_a tree below-ground) represent $\frac{1}{3}$ of the above-ground biomass (Jackson et al. 1996).
285 Given the GPP measured during the dry season was equivalent to GPP of trees (GPP trees) from
286 EC measurements, the carbon use efficiency of the trees (CUE tree) was then calculated (S3, Table
287 S3.1). The resulting CUE value was assessed to determine whether it approximated the typical
288 value of 0.5, which is often used as a default in ecosystem models (Zhou et al., 2019; 2020).

289 2.3.7. Net annual C budget at the ASPS scale

290 The annual C budget of CO_2 fluxes was estimated for chambers and EC measurements in $Mg\ C-CO_2$
291 ha^{-1} . The chambers CO_2 fluxes budgets were obtained by calculating the annual sum of the net
292 CO_2 flux measurements and then weighting with the tree cover rate (10% for the Sh, 90% for the
293 FS). These annual budgets for the field are considered apparent, as they do not account for the
294 biomass exported from the field after the harvest, the decomposition of which therefore escaped
295 both the chambers and the EC. Additionally, the inputs and the outputs of fecal matter resulting
296 from livestock wandering during the dry season were not quantified and are therefore neglected.
297 The objective here is to compare two approaches at different scales using apparent net C budgets,
298 rather than to provide an absolute C budget.



299 *2.4. Statistical analyses*

300 Statistical analyses were performed using the R software (R. Core Team, 2023). To compare the
301 mean values of climatic parameters between the FS and Sh situations, a non-parametric Mann-
302 Whitney test was used when both the normality (shapiro.test) and the homogeneity of the
303 variance (Levene Test, R package ‘Car’; Fox et al., 2023) were not confirmed. This approach was
304 similarly applied to compare the seasonal dynamics of CO₂ fluxes between FS and Sh, as well as
305 between the chamber-based and Eddy Covariance (EC) methods. Means and standard deviations
306 were computed using the ‘skim’ function from the R package ‘skimr’ (Waring et al., 2022).

307 Respiration (R_{ch}) (Eq. 4) and GPP (GPP_{ch}) models (Eq. 7) were fitted using non-linear least
308 squares regression, implemented in the library in R ‘nls.multstart’ (Padfield et al., 2025). For the
309 GPP_{ch} model, parameters α and β with non-significant p-values were removed, and then the
310 remaining values were interpolated and smoothed using a ‘spline’ function from the ‘zoo’ library
311 in R (Zeileis et al., 2024). Ordinary least-square linear regressions were fitted between the
312 measured and the modeled values derived from. Model performance of Eq. 4 and Eq. 7 was
313 evaluated by fitting ordinary least-square linear regressions between the measured and the
314 modeled values using R², root mean square error (RMSE), and the bias metrics. Given that the
315 primary objective of these equations was to accurately reproduce the seasonal dynamics of the
316 CO₂ fluxes to fill gaps in data, particular emphasis was placed on R², with a higher value reflecting
317 a better fit of the model to the measurements.

318 Correlation analysis was conducted between chamber CO₂ fluxes (FCO_{2ch}, R_{ch}, GPP_{ch}) and soil
319 temperature (T_{soil}, °C), air temperature (T_{air}, °C), VWC, the leaf area index of groundnut plants in
320 the chambers (LA_{ch}), and the fitted parameters for respiration — R_{ref} — and photosynthesis —
321 α and β . This analysis was performed using the ‘cor.test’ function from the ‘stats’ package in R
322 (Lüdecke et al., 2021), applying the Spearman method.

323 The threshold of the daily mean soil temperature (T_{soil}, °C) at which the cumulative daily
324 respiration (R_{ch}, g C-CO₂ m⁻² d⁻¹) began to decline was determined using segmented regression
325 from the R package ‘segmented’ (Muggeo, 2003). The associated uncertainty (standard error) of
326 this estimate was evaluated through a bootstrap procedure.

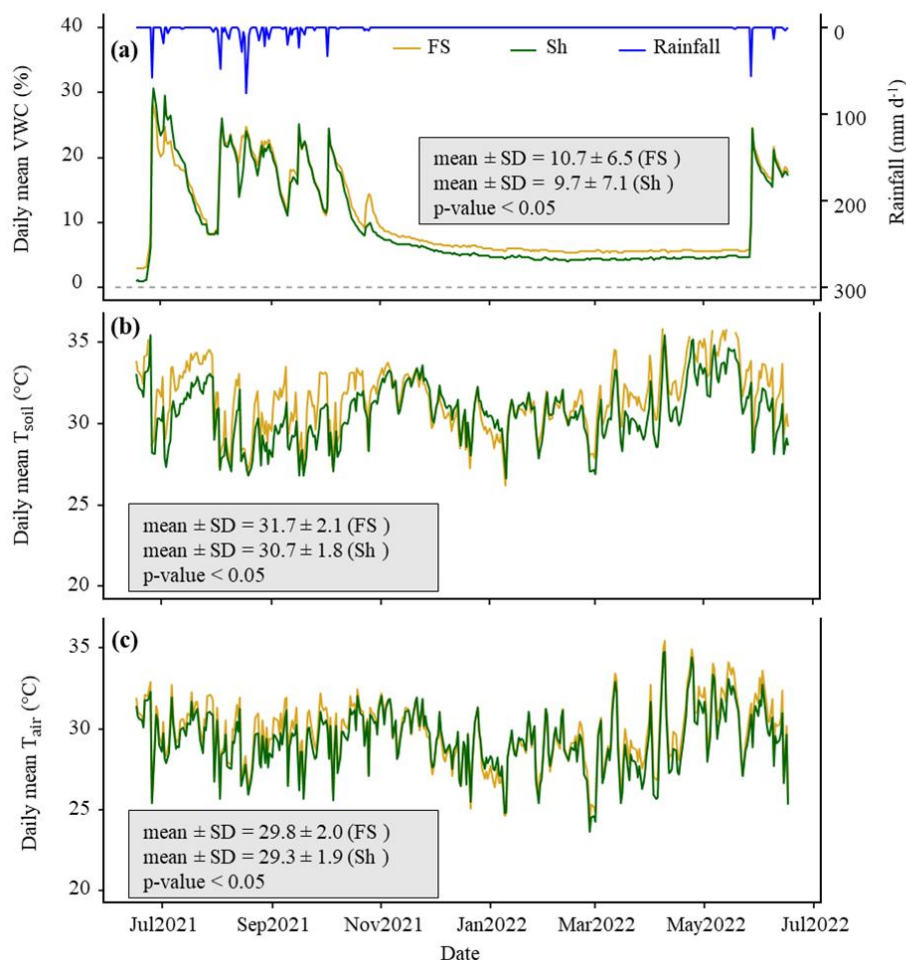


327 3. Results

328 3.1. Microclimatic conditions

329 During the experiment, the cumulative rainfall was 550 mm, which was representative of the
330 interannual average. Precipitations were lowest in July and highest between August and
331 September, a period that typically corresponds to the peak of the rainy season (Fig. 2a). Global
332 radiation ranged between 5.8 and 32.4 MJ m⁻² d⁻¹ (data not shown). The daily mean VWC in the
333 chambers showed significant variation, ranging from 1% at the end of the dry season to a
334 maximum of 30% during the rainy season (Fig. 2a). While VWC was similar during the rainy
335 season, it remained consistently higher in FS than in Sh throughout the dry season ($p < 0.05$),
336 which was unexpected. However, it should be noted that the last rain of October 2021 recharged
337 the FS chambers more effectively, likely due to foliage rainfall interception by *F. albida* which had
338 just put on leaves at that time, potentially explaining this discrepancy in VWC.

339 Within the chamber, the daily mean T_{soil} ranged from 26°C in April to 37.5°C at the end of the dry
340 season (Fig. 2b), while T_{air} varied between 23.7°C and 35.5°C (Fig. 2c). However, during
341 instantaneous daily peaks, T_{soil} could exceed 45°C in May (data not shown). As expected, both daily
342 mean T_{soil} and T_{air} were significantly higher in FS compared to Sh situations ($p < 0.05$), with T_{soil}
343 and T_{air} averaging respectively 1°C and 0.5°C lower under the tree canopy.



344 Fig. 2: One-year time series of daily average microclimatic parameters measured inside chambers.

345 (a) volumetric soil water content (VWC) at a depth of 6 cm (%). (b) soil temperature (T_{soil}) at a depth of 6
346 cm ($^{\circ}\text{C}$), (c) air temperature (T_{air}) at a height of 15 cm ($^{\circ}\text{C}$). The blue line depicts the daily rainfall (mm d^{-1})
347 throughout the year. FS: Full sun chambers; Sh: Shaded chambers. Mean and SD represent respectively the
348 mean value and the standard deviation. The p-value indicates the probability associated with the statistical
349 test, assessing the differences in means between FS and Sh with the significance level α set to 0.05.



350 *3.2. Modeling the chamber-based total respiration (R_{ch}) and photosynthesis (GPP_{ch})*

351 *3.2.1. Dynamics of references respiration, light use efficiency, and maximum CO_2 uptake rate at*
352 *light saturation (R_{ref} , α , and β)*

353 The reference respiration (R_{ref}) showed comparable seasonal dynamics both at a distance from
354 the trees (FS) and under the tree canopies (Sh) (S2, Fig. S2.2). In both situations, R_{ref} showed
355 strong variability during the rainy season, peaking in September 2021 at $2.4 \mu\text{mol } CO_2 \text{ m}^{-2} \text{ s}^{-1}$ for
356 FS and $2.9 \mu\text{mol } CO_2 \text{ m}^{-2} \text{ s}^{-1}$ for Sh (S2, Table S2.1). In contrast, during the dry season — from
357 November 3, 2021 (after harvest) until the onset of the following rainy season (June 2022) — R_{ref}
358 values dropped both for FS and Sh, averaging $0.3 \pm 0.5 \mu\text{mol } CO_2 \text{ m}^{-2} \text{ s}^{-1}$ for FS and $0.5 \pm 0.6 \mu\text{mol}$
359 $CO_2 \text{ m}^{-2} \text{ s}^{-1}$ for Sh. This represents a reduction by a factor of 8 for FS and 6 for Sh compared to the
360 rainy season. The mean annual R_{ref} values were significantly higher under Sh than in FS, with value
361 approximately 1.5 times greater (S2, Table S2.1).

362 Regarding GPP in chambers, the light use efficiency (α) and the maximum CO_2 uptake by
363 groundnut plants in the chambers (β), also reached their maximum during the peak of the rainy
364 season (S2, Fig. S2.3, a and b). The maximum value of α reached $0.2 \mu\text{mol } CO_2 \text{ J}^{-1}$ in FS and 0.3
365 $\mu\text{mol } CO_2 \text{ J}^{-1}$ in Sh (S2, Table S2.1). Similarly, the maximum values of optimum CO_2 uptake rate at
366 light saturation (β) were $40.2 \mu\text{mol } CO_2 \text{ m}^{-2} \text{ s}^{-1}$ for FS and $42.8 \mu\text{mol } CO_2 \text{ m}^{-2} \text{ s}^{-1}$ for Sh (S2, Table
367 S2.1). In the dry season, when photosynthetic activity ceased in the chambers, both α and β were
368 assumed to be nil (S2, Fig. S2.3, a and b). On average, α and β were significantly higher in Sh than
369 in FS, by a factor of 1.7 and 1.2, respectively (S2, Table S2.1). We noted that the decline in
370 photosynthetic activity of the groundnut crop occurred earlier and rapidly at a distance from the
371 trees (FS), as reflected by the sharply observed recession of α and β in FS.

372 *3.2.2. Dynamics of nocturnal respiration in chambers*

373 The averaged nocturnal respiration (nocturnal R_{ch}) calculated from the measurements across
374 each treatment (FS and Sh), showed similar seasonal patterns (Fig. 3, a and c). Following the first
375 rains, R_{ch} values increased dramatically, with a nocturnal 'Birch effect' — a sudden pulse of CO_2
376 release following soil rewetting — observed to be more pronounced under Sh compared to FS,
377 approximately by a factor of 2. At the peak of the rainy season (September), the maximum
378 nocturnal R_{ch} values reached approximately $6.0 \mu\text{mol } CO_2 \text{ m}^{-2} \text{ s}^{-1}$ in FS and $9.0 \mu\text{mol } CO_2 \text{ m}^{-2} \text{ s}^{-1}$
379 in Sh (Fig. 3, a and c). Thereafter, nocturnal R_{ch} declined well before the groundnut harvest along
380 with the rainfall spacing and the groundnut crop senescence (data not shown). During the dry
381 season nocturnal R_{ch} continued to decrease, with maximum values around $1.0 \mu\text{mol } CO_2 \text{ m}^{-2} \text{ s}^{-1}$
382 in FS and $2.0 \mu\text{mol } CO_2 \text{ m}^{-2} \text{ s}^{-1}$ in Sh (Fig. 3, a and c).



383 The modeled nocturnal Rch values closely matched the measured nocturnal Rch values (mean
384 across four chambers per treatment), as indicated by the model performance metrics ($R^2 = 0.9$,
385 with bias and RMSE values of 0.3 and 0.5 $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$, respectively, for FS; $R^2 = 0.7$, with bias
386 and RMSE values of 0.4 and 0.6 $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$, respectively, for Sh) (Fig. 3, b and d). Similarly,
387 the daily mean modeled values also fitted well with the measured values, with FS showing $0.9 \pm$
388 $0.9 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ (modeled) and $1.2 \pm 1.2 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ (measured), while Sh recorded
389 $1.4 \pm 0.9 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ (modeled) and $1.5 \pm 1.2 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ (measured). Given the close
390 match between the measured and modeled values, the fitted model parameters were used
391 subsequently to fill data gaps and estimate diurnal Rch values, as presented in Fig. 4, a and c.

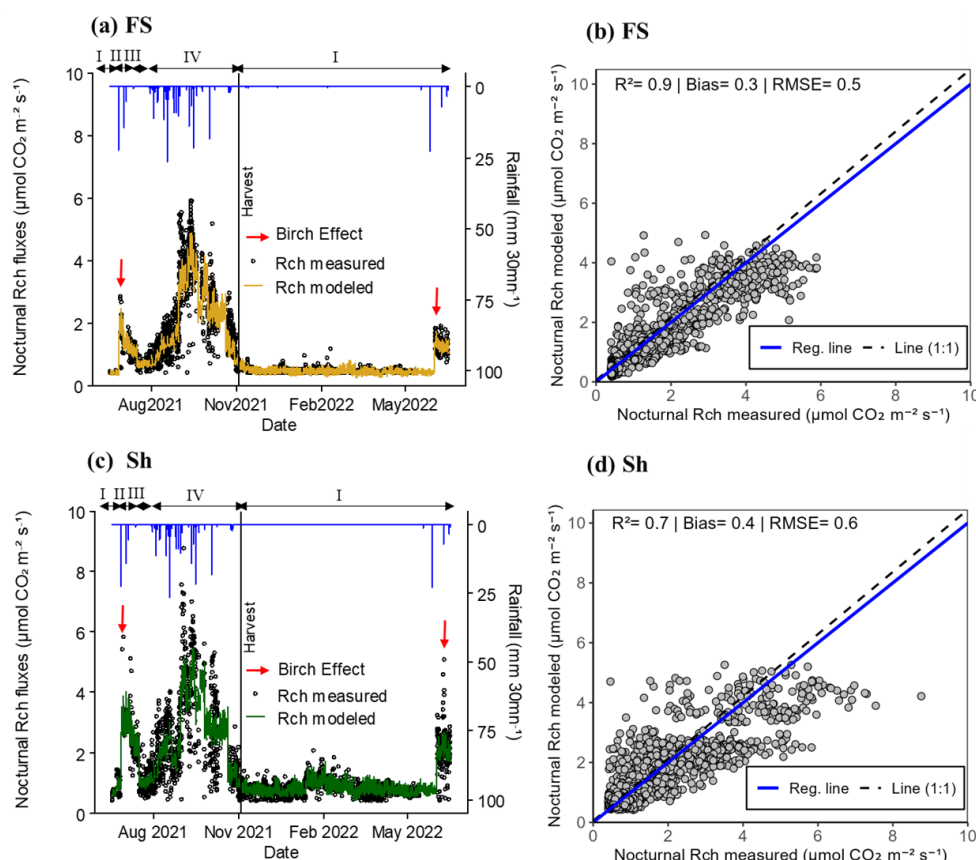


Fig. 3: Dynamics of instantaneous nocturnal CO₂ fluxes in chambers in Full sun (FS; a and b) and Shaded (Sh; c and d) environments (data filtered based on R² of the CO₂ variation over the time of chamber closure and Minimum Detectable Flux, Eq.2).

(a) and (c): measured nocturnal respiration in chambers (Rch: black dots; average of measurements in 4 chambers per location) vs. modeled (coloured line). The vertical black line indicates the harvest date of groundnuts inside the chambers. The red arrows indicate the 'Birch' effect and the blue line represents the rainfall (mm 30mn⁻¹). Roman numerals (above the black arrows) refer to vegetation conditions prevailing inside the chambers, i.e. (I) bare soil, (II) weeds, (III) weeds + groundnuts, and (IV) groundnuts only.

(b) and (d): scatter plot between measured and modeled nocturnal Rch. The solid blue line indicates the regression line and the dashed black one the (1:1) line RMSE and bias are expressed as fluxes (in μmol CO₂ m⁻² s⁻¹). Each point represents the mean value from 4 chambers within the FS or Sh environments.

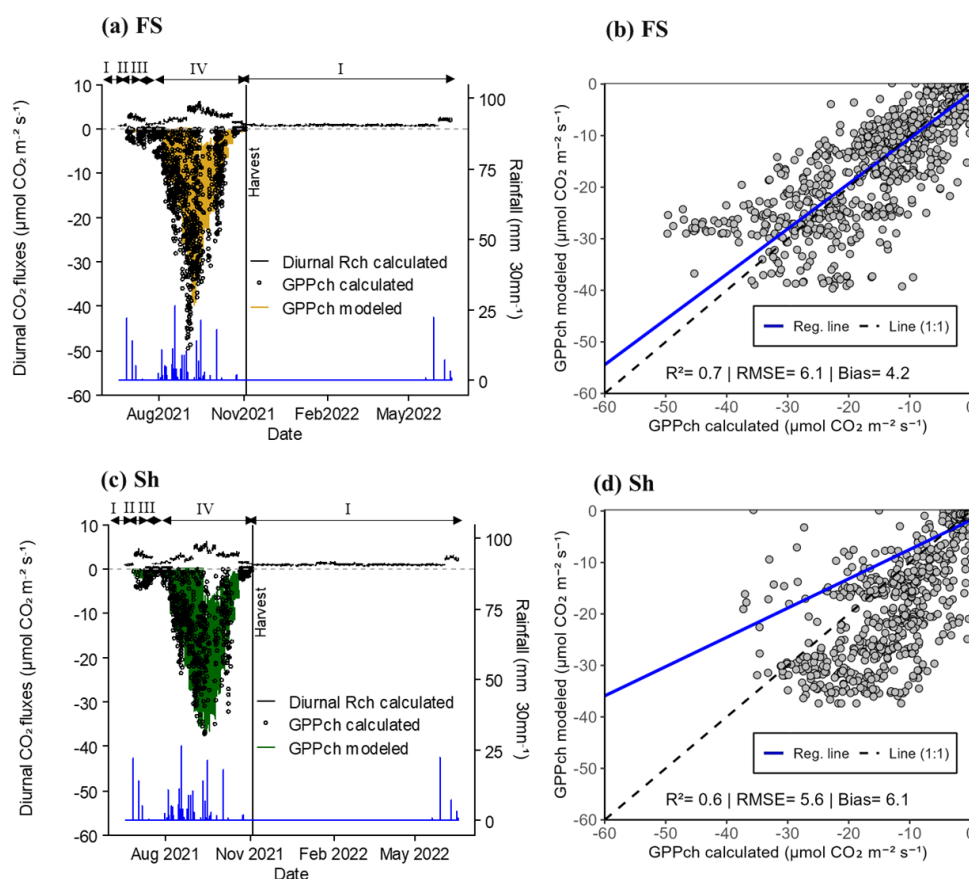


403 3.2.3. *Dynamics of daytime fluxes in chambers*

404 The measured GPP_{ch,stand}, as well as GPP modeled with Eq. 6, showed similar seasonal dynamics
405 in FS and Sh (Fig. 4, a and c). The fluxes peaked during the rainy season (Fig. 4a and c), coinciding
406 with periods of vigorous vegetative growth characterised by a high leaf area index (LAI_{ch}) of
407 groundnut plants within the chambers (S2, Fig. S2.1). The maximum calculated and standardised
408 GPP_{ch} values reached $-50 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ for FS and $-37 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ for Sh. As expected,
409 these fluxes were nil during the dry season when the soil was bare (Fig. 4, a and c).

410 The modeled GPP_{ch} values closely followed the same trends as the calculated values, although
411 model performance was slightly better for FS ($R^2 = 0.7$ with bias and RMSE values of 4.2 and 6.1
412 $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$, respectively) compared to Sh ($R^2 = 0.6$ with bias and RMSE values of 6.1 and
413 5.6 $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$, respectively) (Fig. 4, b and d).

414 The calculated diurnal respiration values (diurnal R_{ch} calculated) for FS and Sh revealed a 'Birch
415 effect' similar to that observed during the night, though slightly more pronounced under Sh by a
416 factor of 1.2. Diurnal R_{ch} values increased significantly during the rainy season, reaching a
417 maximum of $6.0 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ for both FS and Sh (Fig. 4, a and c). In the dry season, on bare
418 soil, these values declined, with maximum respiration reaching only $0.5 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ for both
419 situations (FS and Sh) (Fig. 4, a and c).



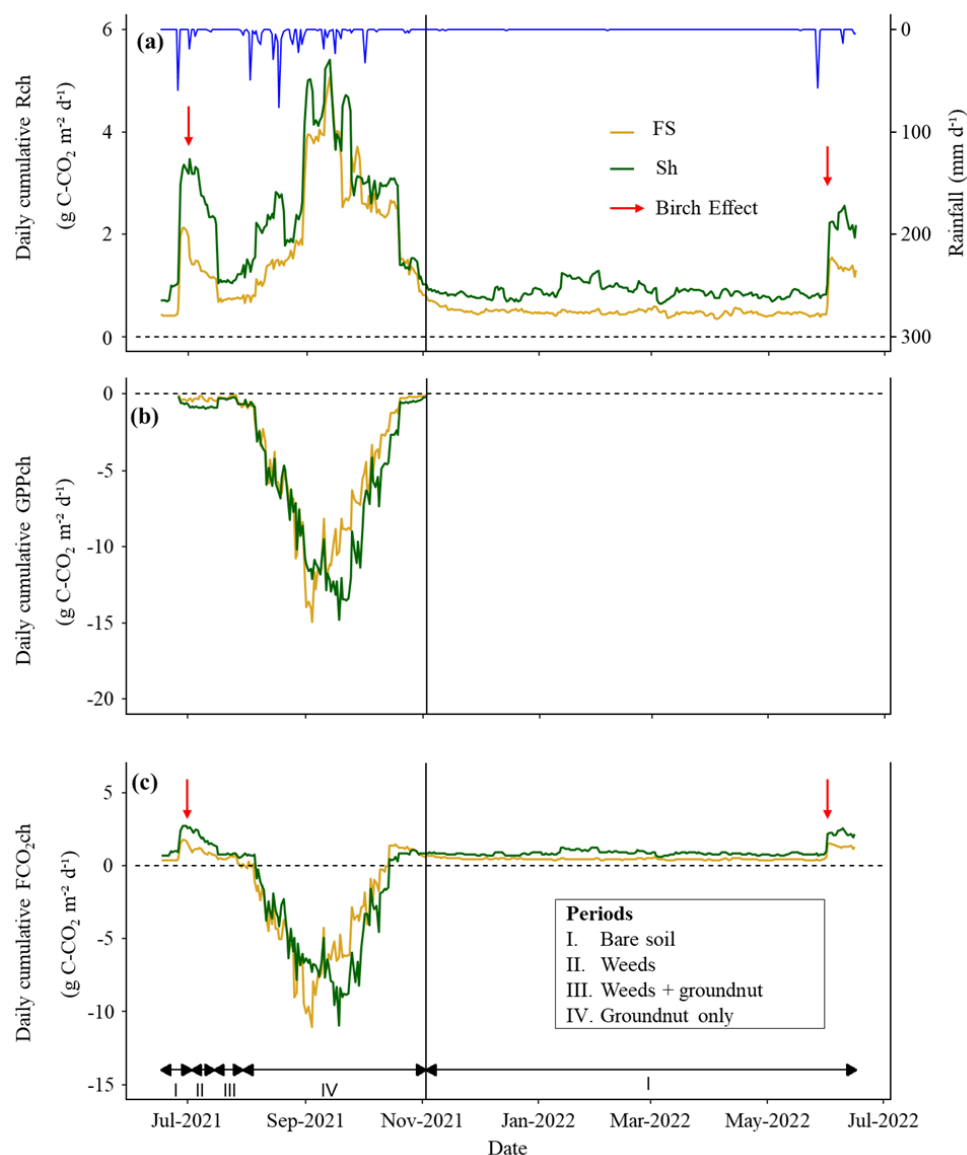
420 Fig. 4: Dynamics of instantaneous diurnal CO₂ fluxes in chambers in Full sun (FS; a and b) and
421 Shaded (Sh; b and d) environments (filtered based on R² of the CO₂ variation over the time closure
422 in FS and Sh and Minimum Detectable Flux, Eq.2).

423 (a) and (c): non-gap-filled diurnal Rch calculated (black line, positive values; average of measurements in
424 4 chambers per location) and GPPch calculated from Eq.5 then standardised for LAI (black dots, negative
425 values) and modeled (coloured line, negative values). The vertical black line indicates the harvest date of
426 groundnuts inside the chambers and the blue line represents the rainfall (mm 30mn⁻¹). Roman numerals
427 (above the black arrows) refer to conditions prevailing inside the chambers, i.e., (I) bare soil, (II) weeds,
428 (III) weeds + groundnuts, and (IV) groundnuts.
429 (b) and (d): scatter plot between calculated and modeled GPPch. The solid blue line indicates the regression
430 line and the dashed black one the (1:1) line. RMSE and bias are expressed as fluxes (in μmol CO₂ m⁻² s⁻¹).
431 Each point represents the mean value from 4 chambers within the FS or Sh environments.



432 *3.3. Dynamics of daily cumulative CO₂ fluxes in chambers*

433 The seasonality of daily cumulative of GPPch.stand showed similar dynamics between FS and Sh,
434 with higher variability during the rainy season than during the dry season (Fig. 5). Daily total Rch
435 peaked during the rainy season at 5.1 g C-CO₂ m⁻² d⁻¹ for FS and 5.4 g C-CO₂ m⁻² d⁻¹ for Sh, while
436 the maximum GPPch.stand values were comparable at around -15.0 g C-CO₂ m⁻² d⁻¹ for both FS
437 and Sh (Table 1; S2, Fig. S2.4, a, b, c, and d). In the dry season, Rch decreased (Fig. 5), averaging
438 0.5 g C-CO₂ m⁻² d⁻¹ for FS and 1.0 g C-CO₂ m⁻² d⁻¹ for Sh. GPPch declined well before harvest
439 (senescence) and remained nil during the dry season (Fig. 5). During the rainy season FCO₂ch
440 peaked at around 11.0 g C-CO₂ m⁻² d⁻¹ for FS and Sh (Fig. 5) (Table 1; S2, Fig. S2.4, e and f), while
441 FCO₂ch values were the same as Rch during the dry season. In absolute terms, the mean Rch and
442 GPPch were significantly higher under Sh as compared to FS, by factors of 1.3 and 1.2, respectively.
443 Conversely, the mean FCO₂ch was significantly higher under FS (0.4 g C-CO₂ m⁻² d⁻¹) than under
444 Sh (0.2 g C-CO₂ m⁻² d⁻¹) (Table 1).
445 The annual cumulative Rch values were 392.8 g C-CO₂ m⁻² for FS and 574.5 g C-CO₂ m⁻² for Sh.
446 The GPPch fluxes reached -539.5 g C-CO₂ m⁻² for FS and -632.6 g C-CO₂ m⁻² for Sh. Annual net
447 cumulative C exchange (FCO₂ch) were -146.7 g C-CO₂ m⁻² in FS and -58.1 g C-CO₂ m⁻² in Sh.



448 Fig. 5: Seasonal dynamics of daily gap-filled cumulative fluxes (in gC-CO₂ m⁻² d⁻¹) in chambers.

449 (a) soil+crop respiration (Rch), (b) photosynthesis (GPPch, standardised for LAI) and (c) net CO₂ exchange
 450 (FCO₂ch). The yellow and green solid lines compare the FS and Sh environments, respectively. The vertical
 451 black line indicates the harvest date of groundnuts inside the chambers. The blue line depicts the daily
 452 cumulative rainfall (mm d⁻¹) throughout the rainy season, and the red arrow indicates the 'Birch'
 453 effect. Roman numerals (above the black arrows) in (a) and (c) refer to the prevailing conditions inside the
 454 chambers: (I) bare soil, (II) weeds, (III) weeds + groundnuts, (IV) groundnuts.

455



Table 1: Comparison of daily cumulative and gap-filled chamber CO₂ fluxes (R_{ch}, GPP_{ch} standardised for LAI, and FCO_{2ch} in g C-CO₂ m⁻²) in the FS and Sh condition.

	Annual sum	Daily Mean \pm SD	Min	Max	Mann-Whitney test
(g C-CO ₂ m ⁻²)	.yr ⁻¹	.d ⁻¹	.d ⁻¹	.d ⁻¹	
Rch					
FS	392.8	1.1 \pm 0.9	0.4	5.1	*
Sh	574.5	1.6 \pm 1.1	0.6	5.4	
GPPch					
FS	-539.5	-4.1 \pm 4.3	< -0.1	-14.9	*
Sh	-632.6	-4.8 \pm 4.6	< -0.1	-14.8	
FCO₂ch					
FS	-146.7	-0.4 \pm 2.4	-11.0	1.8	*
Sh	-58.1	-0.2 \pm 2.7	-10.9	2.8	

Annual sum corresponds to the annual cumulative fluxes (g C-CO₂ m⁻² yr⁻¹). Mean, SD, Min, and Max represent respectively the mean, standard deviation, minimum, and maximum values at the daily scale (g C-CO₂ m⁻² d⁻¹). Asterisks (*) indicate the p-values from the Mann-Whitney test, used to assess differences in mean between FS and Sh (p < 0.05). Positive values indicate CO₂ emissions, while negative values represent CO₂ uptake.



464 *3.4. Drivers of daily respiration and photosynthesis in chambers*

465 The chamber-based daily cumulative respiration (R_{ch}) and GPP_{ch} showed significant and positive
466 correlations with the leaf area index (LAI_{ch}), both at a distance from the trees (FS) and under the
467 trees (Sh) (Table 2). The influence of LAI_{ch} on GPP_{ch} was stronger ($r = 0.86$ for FS and Sh) than
468 its influence on R_{ch} ($r = 0.60$ for FS; $r = 0.69$ for Sh). Soil VWC was also positively correlated with
469 R_{ch} and GPP_{ch}, both in FS and Sh. However, the influence of soil VWC on R_{ch} was stronger under
470 Sh compared to FS, while its influence on GPP_{ch} was similar in both situations (FS and Sh). Soil
471 temperature showed weak negative correlations with R_{ch} (in FS and Sh) and with GPP_{ch} (only in
472 Sh). Finally, no significant correlations were found between T_{air} and any of the CO₂ fluxes (Table
473 2).



474 Table 2: Spearman correlation matrix based on daily cumulative and gap-filled CO₂ fluxes from full
475 year chamber measurements (g C-CO₂ m⁻² d⁻¹) with microclimatic parameters.

Parameters	Condition	Rch	GPPch
T _{soil}	FS	-0.25 ***	ns
	Sh	-0.28 ***	-0.38 ***
T _{air}	FS	ns	ns
	Sh	ns	ns
VWC	FS	0.51 ***	0.75 ***
	Sh	0.78 ***	0.75***
LAIch	FS	0.60 ***	0.86 ***
	Sh	0.69 ***	0.86 ***

476 Spearman correlation coefficients between daily cumulative and gap-filled CO₂ flux components (Rch and
477 GPPch, with GPPch in absolute terms) and daily mean microclimatic parameters in full sun (FS) and shaded
478 chambers (Sh). T_{soil} (°C) is the daily mean soil temperature at 6 cm depth, T_{air} (°C) the daily mean air
479 temperature at 15 cm height, VWC (%) the daily mean volumetric water content (VWC, %), and LAIch (m⁻²
480 leaf m⁻² soil) the chamber leaf area index value for a given day. Significance levels are indicated by (***) for
481 p < 0.001; ns denotes a non-significant correlation (p > 0.05)



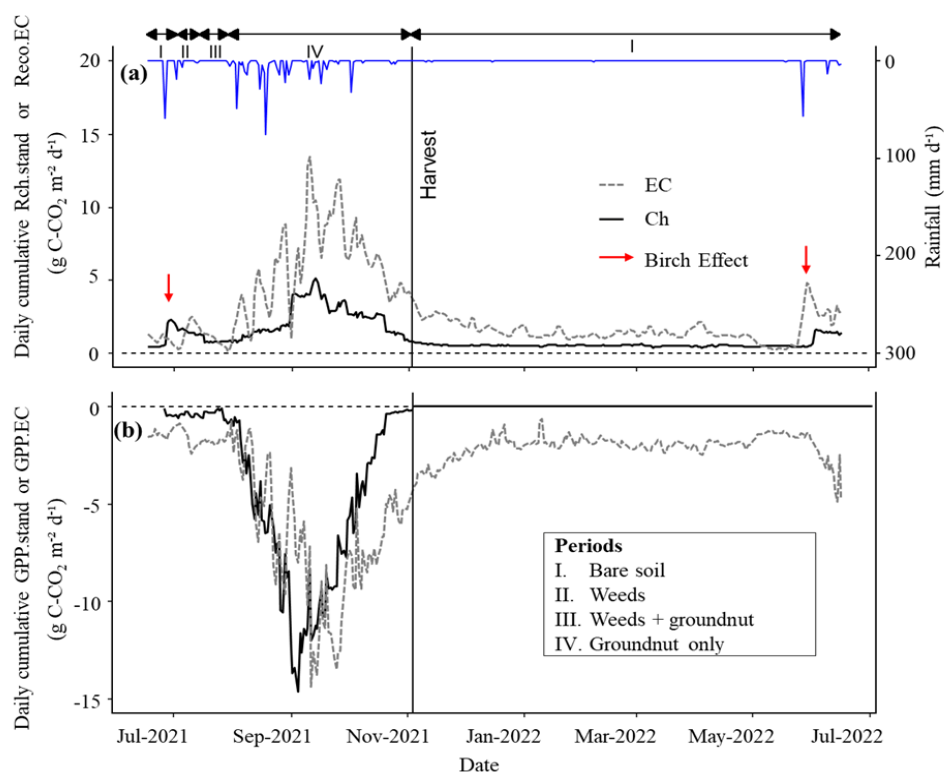
482 *3.5. Comparison of respiration and GPP measurements between chambers (Ch) and Eddy*
483 *Covariance (EC) methods*

484 The chamber-based daily total CO₂ fluxes, gap-filled and weighted according tree cover were
485 compared with the fluxes obtained using the EC method (Fig. 6).

486 During the rainy season, both total respiration and GPP showed comparable dynamics between
487 the two methods, with synchronised peaks and higher variability compared to the dry season (Fig.
488 6). The maximum value of Reco.EC, peaked at 13.5 g C-CO₂ m⁻² d⁻¹ (Table 3). The initial value of
489 Rch.stand was comparable to Reco.EC but peaked only at 5.1 g C-CO₂ m⁻² d⁻¹ (Table 3), meaning
490 a third of the peak of Reco.EC. The maximum GPP, was -14.3 g C-CO₂ m⁻² d⁻¹ and -14.6 g C-CO₂
491 m⁻² d⁻¹ for GPP.EC and GPPch.stand, respectively (Table 3). This indicates that the LAI-based
492 standardisation and upscaling approach were realistic, at least up to the peak of groundnut
493 growth.

494 On average, Reco.EC was significantly higher than Rch.stand, by a factor of 2.3. GPP.EC was also
495 significantly higher than GPPch.stand, but only by a factor of 1.2 (Table 3).

496 During the dry season, Reco.EC and Rch.stand gradually decreased. The values for Reco.EC
497 remained higher than for Rch.stand, which was fairly consistent with the contribution of the Ra
498 tree above-ground compartment, even if this difference seemed to disappear at the end of the dry
499 season (Fig. 6). The measured 'Birch effect' was highest for Rch.stand in 2021, but was the
500 opposite in 2022 due to a system failure at the beginning of the rainy season. The maximum value
501 of GPP.EC reached -2.4 g C-CO₂ m⁻² d⁻¹ when the trees were at their maximum of foliage, after
502 harvest and while weeds were still present in the field. However, after the harvest, chamber
503 photosynthesis (GPPch.stand) was nil (Table 3).



504 Fig 6: Comparing the seasonal dynamics of CO₂ fluxes between Eddy Covariance (EC)
505 measurements and upscaled chamber measurements (ch.stand).

506 (a) represent the seasonal dynamics of soil + crop respiration (Rch.stand) and ecosystem respiration
507 (Reco.EC) and (b) photosynthesis (GPP.stand and GPP.EC). The black and dashed grey lines show Ch and
508 EC seasonal dynamics, respectively. The vertical black line indicates the harvest date of groundnuts inside
509 the chambers. The blue line depicts the daily cumulative rainfall (mm d⁻¹), and the red arrow indicates the
510 'Birch' effect. Roman numerals (above the black arrows) refer to conditions prevailing inside the
511 chambers: (I) bare soil, (II) weeds, (III) weeds + groundnuts, (IV) groundnuts.

Table 3: Comparison of gap-filled CO₂ fluxes between Eddy Covariance (EC) and upscaled chamber (Ch.stand) measurements, by season (rainy or dry).

	Rainy season					Dry season				
	Daily Mean ± SD .d ⁻¹	Min .d ⁻¹	Max .d ⁻¹	Mann-Whitney test		Daily Mean ± SD .d ⁻¹	Min .d ⁻¹	Max .d ⁻¹	Mann-Whitney test	
Reco.EC or Rch.stand										
EC	4.6 ± 3.2	0.2	13.5			1.2 ± 0.4	0.3	2.1		*
Ch.stand	2.0 ± 1.1	0.5	5.1			0.5 ± 0.04	0.4	0.6		
GPP.EC or GPPch.stand										
EC	-5.1 ± 3.6	-0.7	-14.3		*	-1.7 ± 0.3	-0.6	-2.4		
Ch.stand	-4.2 ± 4.3	<-0.1	-14.6			0	0	0		-

Mean, SD, Min, and Max represent the daily mean fluxes, standard deviation, minimum, and maximum values, respectively (g C-CO₂ m⁻² d⁻¹). The Asterisks (*) indicate the p-values from the Mann-Whitney test, used to assess differences in mean between EC and Ch. Positive values indicate CO₂ emissions, while negative values represent CO₂ uptake.



516 3.6. *The contribution of *F. albida* to Reco and GPP*

517 During the dry season, the cumulative contribution of *F. albida* to ecosystem respiration (Ra tree)
518 was 139.6 g C-CO₂ m⁻². This represent 12% of the total annual cumulative Reco, which was
519 estimated at 1180.0 g C-CO₂ m⁻². The contribution of trees (GPP tree) to total annual GPP was -
520 270.2 g C-CO₂ m⁻², equivalent to ~50% of the total annual cumulative GPP of the ecosystem (550
521 g C-CO₂ m⁻²).
522 The ratio between these two components (Ra tree / GPP tree) in absolute terms was 0.52,
523 reflecting a carbon use efficiency (CUE) of 0.48 (S3, Table S3.1).

524 3.7. *Carbon budgets at the field-scale*

525 The upscaled chamber-based annual cumulative total respiration flux (Rch.stand) was estimated
526 to be 4.1 ± 0.01 Mg C-CO₂ ha⁻¹ (Table 4). In comparison, the annual budget of Reco.EC was 10.0
527 ± 0.03 Mg C-CO₂ ha⁻¹ (Table 4), more than two times larger than Rch.stand.
528 The upscaled GPPch.stand reached an annual cumulative value of -5.5 ± 0.03 Mg C-CO₂ ha⁻¹,
529 whereas the annual cumulative GPP.EC was -11.8 ± 0.03 Mg C-CO₂ ha⁻¹ (Table 4).
530 The annual net C budget, based on both methods, was estimated at -1.4 ± 0.02 Mg C-CO₂ ha⁻¹ for
531 chambers (FCO₂ch.stand) and -1.8 ± 0.01 Mg C-CO₂ ha⁻¹ for Eddy Covariance (NEE.EC) (Table 4).



532 Table 4: Annual budget of CO₂ fluxes based on Eddy Covariance (EC) and upscaled chamber
533 methods (Ch.stand).

	Annual sum	Std error
	(Mg C-CO ₂ ha ⁻¹)	(Mg C-CO ₂ ha ⁻¹)
Reco.EC or Rch.stand		
EC	10.0	0.03
Ch.stand	4.1	0.01
GPP.EC or GPPch.stand		
EC	-11.8	0.03
Ch.stand	-5.5	0.03
NEE.EC or FCO₂ch.stand		
EC	-1.8	0.01
Ch.stand	-1.4	0.02

534 Annual sum corresponds to the annual cumulative fluxes for full year measurements (Mg C-CO₂ ha⁻¹). EC
535 refers to fluxes measured by the Eddy Covariance method, and Ch refers to the fluxes measured by
536 chambers, which are then upscaled to the whole field. Rch.stand represents the chamber respiration, while
537 Reco.EC denotes the ecosystem respiration according to the EC method. GPP.EC and GPPch.stand are the
538 gross primary production or photosynthesis flux, measured by EC and Ch methods, respectively. NEE.EC
539 and FCO₂ch.stand represent the net ecosystem exchange for EC and Ch, respectively. The associated
540 standard error is denoted as Std error (Mg C-CO₂ ha⁻¹). Positive values indicate CO₂ emissions, while
541 negative values represent CO₂ uptake.



542 4. Discussion

543 4.1. Seasonality and drivers of chamber-based CO₂ fluxes

544 In our agroforestry context, seasonal variability in CO₂ fluxes closely followed rainfall dynamics,
545 peaking during the wet season and declining sharply in the dry season, consistent with soil
546 moisture depletion and crop senescence. This pattern is typical of semi-arid ecosystems (Ago et
547 al., 2016a; Brümmer et al., 2008; Guillen-Cruz et al., 2023; Macharia et al., 2020; Mosongo et al.,
548 2022; Wieckowski et al., 2024).

549 Respiration and photosynthesis were primarily driven by soil moisture and LAI, reflecting the
550 system's sensitivity to water availability and crop dynamics. Soil moisture enhanced both
551 processes by stimulating microbial activity and supporting plant growth (Borken et al., 2002;
552 Conant et al., 2004; Merbold et al., 2009; Yu et al., 2020; Zhao et al., 2016). The stronger correlation
553 between soil moisture and respiration under F. albida canopy (Sh: $r = 0.78$) compared to full sun
554 (FS: $r = 0.51$) suggests greater microbial sensitivity to moisture beneath trees. This likely reflects
555 enhanced substrate availability, resulting in stronger post-rainfall respiration pulses (Meisner et
556 al., 2015) and supporting the 'fertile island' effect, where trees improve local soil conditions
557 (Eldridge et al., 2024). Photosynthetic capacity also responded to soil moisture, as shown by
558 positive correlations with LAI and key physiological traits such as light use efficiency (α) and
559 maximum CO₂ uptake rate (β) (Gonsamo et al., 2019; Qiu et al., 2023; Zhang et al., 2024).

560 In contrast, the influence of soil temperature (T_{soil}) on respiration was weakly negative in both FS
561 and Sh, indicating a thermal threshold beyond which respiration is suppressed—estimated at 32
562 ± 1.5 °C in FS and 29.5 ± 1.9 °C in Sh (S2, Fig. S2.6, a and b), similar to findings in Eastern Ghana
563 (Owusu et al., 2024). This inhibition likely results from decreased enzymatic and microbial
564 activity under combined heat and water stress (Liu et al., 2018; Richardson et al., 2012). In semi-
565 arid regions, soil respiration often becomes decoupled from temperature due to seasonal
566 moisture constraints (Jia et al., 2020; Tucker & Reed, 2016; Warren, 2014), with microbial activity
567 limited during dry periods despite favourable temperatures. This decoupling helps explain the
568 weak or absent correlation between T_{soil} and soil moisture (S2, Fig. S2.5, b), particularly under Sh
569 ($r = -0.28$). Management practices such as organic inputs can also modulate these dynamics,
570 adding further variability to soil respiration responses (Meena et al., 2020; Oyonarte et al., 2012;
571 Rong et al., 2015; Xue & Tang, 2018).

572 4.2. Magnitude of chamber-based total CO₂ respiration fluxes

573 Mean total soil respiration values were consistent with those reported in other low-input
574 agricultural systems across sub-Saharan Africa (Mapanda et al., 2010; Pelster et al., 2017;
575 Rosenstock et al., 2016). In full sun (FS), the mean respiration (1.0 ± 0.9 g C-CO₂ m⁻² d⁻¹) closely



576 matched values measured by Wachiye et al. (2020) in a semi-arid Kenyan field at 1158 m altitude
577 ($1.1 \pm 0.1 \text{ g C-CO}_2 \text{ m}^{-2} \text{ d}^{-1}$). This similarity likely reflects comparable environmental conditions,
578 including moderate rainfall ($\sim 550 \text{ mm yr}^{-1}$) and low soil organic carbon and nitrogen contents
579 ($<1\%$) in the 0–20 cm layer of sandy soil. In contrast, respiration under *F. albida* canopy (Sh: 1.6
580 $\pm 1.1 \text{ g C-CO}_2 \text{ m}^{-2} \text{ d}^{-1}$) was higher, likely due to additional autotrophic respiration from tree roots
581 and greater organic inputs beneath the canopy. Nonetheless, this flux remains close to values
582 observed in low-input sorghum fields on sandy loam soils in eastern Ghana ($1.7 \pm 1.1 \text{ g C-CO}_2 \text{ m}^{-2}$
583 d^{-1}), despite higher rainfall ($950\text{--}1000 \text{ mm yr}^{-1}$) in that region (Owusu et al., 2024).
584 Cumulative annual respiration fluxes fell within the range reported for Sahelian croplands (250--
585 $450 \text{ g C-CO}_2 \text{ m}^{-2}$) (Brümmer et al., 2009) and other sub-Saharan African agricultural systems (Kim
586 et al., 2016). The cumulative flux under tree cover is similar to that measured in cassava fields in
587 eastern Tanzania ($440 \text{ g C-CO}_2 \text{ m}^{-2} \text{ yr}^{-1}$), despite the latter receiving higher rainfall ($\sim 1115 \text{ mm}$
588 yr^{-1}) (Rosenstock et al., 2016). This convergence may stem from comparable soil fertility
589 constraints, with low soil organic carbon ($1\text{--}1.7\%$) and nitrogen contents ($<0.5\%$). In contrast,
590 the slightly lower cumulative flux in FS may reflect less favourable microclimatic conditions—
591 such as elevated soil temperatures and increased aridity away from tree cover—limiting
592 microbial activity (see Section 4.1).
593 Across sub-Saharan Africa, soil respiration fluxes based on static chamber measurements show
594 high spatial variability, largely shaped by climate and land use. For example, Owusu et al. (2024)
595 found higher respiration in woodlands ($3.8 \pm 0.8 \text{ g C-CO}_2 \text{ m}^{-2} \text{ d}^{-1}$) and grazed areas (2.7 ± 1.7)
596 than in croplands (1.7 ± 1.1) in humid eastern Ghana. This gradient was linked to differences in
597 soil moisture and organic matter. Similarly, Rosenstock et al. (2016) reported much higher fluxes
598 in highland pastures in Kenya ($3.8\text{--}4.4 \text{ g C-CO}_2 \text{ m}^{-2} \text{ d}^{-1}$) compared to cultivated fields in eastern
599 Tanzania (1.2 ± 0.2), highlighting the role of vegetation cover and soil fertility.

600 4.3. Effect of trees on chamber-based soil respiration and photosynthesis

601 A notable increase in respiration and photosynthesis fluxes was observed under *F. albida* trees
602 (Sh) compared at a distance from trees (FS). This increase may indicate the potential role of *F.*
603 *albida* in modulating CO_2 exchange dynamics (Rch and GPPch) within this agro-silvo-pastoral
604 system. These results are consistent with preliminary findings from similar environments
605 (Duthoit et al., 2020).
606 Numerous studies have investigated the effect of tree species on greenhouse gas fluxes,
607 particularly CO_2 , revealing significant variations across different ecological contexts (Bréchet et
608 al., 2021, 2025; Klaus et al., 2024; Mazza et al., 2021; Ramesh et al., 2013; Rheault et al., 2024).
609 However, the underlying mechanisms by which trees influence these dynamics are not yet fully
610 understood.



611 In general, agroforestry systems have been well-documented for their ability to provide a range
612 of ecosystem services (e.g., Assefa et al., 2024; Bado et al., 2021; Kuyah et al., 2019; Rolo et al.,
613 2023). Specifically, *Faidherbia*-based agroforestry systems may play a crucial role in regulating
614 CO₂ exchanges between the soil and atmosphere. *F. albida*-based agroforestry systems are
615 recognized for enhancing both soil organic and mineral fertility (Bayala et al., 2020; Dilla et al.,
616 2019; Sileshi, 2016; Sileshi et al., 2020; Stephen et al., 2020), mainly through litter accumulation
617 and direct inputs from livestock excreta under their canopies. Additionally, the extensive roots
618 system of *F. albida* trees helps concentrate mineral nutrients, contributing to the formation of a
619 'fertile island' effect under the trees (Siegwart et al., 2022; Eldridge et al., 2024). Moreover, *F.*
620 *albida* improve water infiltration (Diongue et al., 2023; Faye et al., 2020; Sarr et al., 2023), enhance
621 soil moisture retention (Clermont-Dauphin et al., 2023) and contribute to reduced soil
622 temperatures (de Carvalho et al., 2021; Lopes et al., 2024; Sida et al., 2018). These changes foster
623 a more favourable environment for soil microbial activity and crop development (Diack et al.,
624 2024; Diene et al., 2024; Leroux et al., 2020; Rouspard et al., 2020) under the trees compared to
625 open areas. Consequently, this likely explains the stronger effect of soil moisture and the leaf area
626 index of groundnuts on R_{ch} under the trees, resulting in higher total respiration (Table 2). For
627 photosynthesis, the effect of these parameters was similar in both FS and Sh (Table 2). However,
628 the significantly higher intensity of GPP_{ch} under Sh can be explained by greater light use efficiency
629 (α) and a higher maximum CO₂ uptake rate at light saturation (β) in this shaded environment. In
630 agroforestry systems, light use efficiency can at least partially mitigate the reduction in
631 photosynthetically active radiation under tree canopies (Charbonnier et al., 2017).
632 Similar results have been observed in different climatic conditions and ecosystems. Gomes et al.
633 (2016) investigated soil respiration using mobile chambers (LI-8100-102 model) under trees in
634 coffee-based agroforestry (AF) systems and in open areas (FS) in Minas Gerais, Brazil. These
635 studies were conducted with agroecological management practices, such as weeding,
636 intercropping maize between coffee rows, and mulching. The AF systems exhibited lower air and
637 soil temperatures (at 5 and 10 cm depth) and higher air and soil humidity compared to FS (Gomes
638 et al., 2016). These authors observed greater spatial variability in soil respiration in AF (34.1%)
639 compared to FS (24.2%). This variability was mainly linked with fluctuations in labile carbon and
640 total nitrogen, reflecting more favourable soil microclimate for microbial activity in AF. In
641 contrast, soil temperature (10 cm depth) accounted for most of the variability observed in FS,
642 where the absence of tree canopy resulted in high soil temperatures and low soil moisture (Gomes
643 et al., 2016). Likewise, Haren et al. (2010) reported 38% higher soil respiration near large trees
644 (DBH > 35 cm) in clay-rich Amazonian forests compared to open sites. Interestingly, the
645 magnitude of CO₂ fluxes was independent of tree species, indicating that canopy effects may
646 outweigh species-specific traits in some contexts. In our study, *F. albida*'s influence on CO₂ fluxes



aligns with this general pattern observed in tropical agroforestry. However, the mechanisms linking individual tree species to microbial and physicochemical drivers of CO₂ dynamics remain insufficiently understood and warrant further investigation (Jevon et al., 2023).

4.4. Birch Effect

A rapid increase in soil respiration was observed following the first rainfall events, particularly under *F. albida*. This phenomenon can be attributed to the lower bulk density of the soil under the trees (Clermont-Dauphin et al., 2023; Siegwart et al., 2023), which potentially lead to CO₂ accumulation during the dry season due to higher soil organic matter (SOM) (Siegwart et al., 2023). Additionally, the sensitivity of microbial communities to subtle variations in soil moisture, compounded by the tree effect, may further explain this phenomenon, as outlined in Sections 4.1 and 4.3. This phenomenon, known as the 'Birch effect' (Birch, 1958), has been reported across various semi-arid ecosystems in sub-Saharan Africa (Ago et al., 2016b; Fan et al., 2015; Wieckowski et al., 2024), as well as other semi-arid ecosystems globally (Roby et al., 2022; Yan et al., 2014; Yu et al., 2020). In these contexts, the 'Birch effect' may result from the displacement of soil gas phases by the piston effect generated during water infiltration (Singh et al., 2023). Furthermore, microbial communities in semi-arid environments adopt osmoregulatory mechanisms to withstand water deficit (Warren, 2014), which is particularly pronounced during the dry season. This phenomenon reduces soil microbial metabolism (Schimel et al., 2007). Upon rapid soil rewetting, especially after prolonged dry periods, soil microbial metabolism process is swiftly reactivated, leading to a transient pulse in respiration and a CO₂ release (Barnard et al., 2020; Kim et al., 2012; Manzoni et al., 2020; Vargas et al., 2018). Isotopic signatures of soil respiration provide evidence supporting the hypothesis that these pulses result from the rapid mineralisation of necromass or osmolytes excreted by microorganisms under drought stress (Schimel et al., 2007; Unger et al., 2010). Additional factors may amplify the 'Birch effect'. For instance, drying-rewetting cycles can induce physical disruption of soil aggregates, enhance oxygen penetration and thereby expose previously protected organic matter to microbial decomposition (Rabbi et al., 2024). This increases substrate availability and subsequently boosts soil respiration fluxes.

The magnitude of the 'Birch effect' is modulated by the severity and duration of drought. Thus, at our study site, given the 8- to 9-month-long dry season, the 'Birch effect' is particularly intense. Indeed, extended drought periods promote greater accumulation of microbial necromass and intensify hypo-osmotic stress responses upon rewetting (Singh et al., 2023).



679 *4.5. Comparing EC and chamber-based methods*

680 Results revealed high seasonal variability, with higher values during the rainy season compared
681 to the dry season. This seasonal pattern aligns with findings from studies in the Sahel using the
682 EC method for CO₂ flux measurements (Brümmer et al., 2008; Tagesson et al., 2015; Agbohessou
683 et al., 2023, Wieckowski et al., 2024). Comparable patterns have been also documented at the
684 ecosystem scale in other semi-arid environments (Ago et al., 2014; Archibald et al., 2009; Ardö et
685 al., 2008; Jia et al., 2020; Quansah et al., 2015; Williams et al., 2009; Zhang, Bi, et al., 2024).

686 Several comparative studies between chamber and EC methods have reported both congruent
687 and divergent CO₂ flux estimates (Bastviken et al., 2022; Poyda et al., 2017; Riederer et al., 2014;
688 J. Tang et al., 2008; Wang et al., 2010). In the present study, ecosystem respiration fluxes during
689 the rainy season exhibited notable discrepancies measurements between EC (Reco.EC) and
690 upscaled chamber-based (Rch.stand). This is attributable to differences in the flux components
691 captured by each method. Specifically, Reco.EC included respiration from below- and above-
692 ground tree parts, crops (groundnuts and cowpeas), weeds, and soil, whereas Rch.stand
693 accounted only respiration from below-ground tree, groundnut crop, and soil. Therefore, as
694 expected, Reco.EC (4.6 ± 3.2 g C-CO₂ m⁻² d⁻¹) were significantly higher than Rch.stand (2.0 ± 1.1 g
695 C-CO₂ m⁻² d⁻¹).

696 For chamber-based GPP measurements, values were standardised (GPP-stand) by the field's leaf
697 area index (LAI.field). This allowed it to improve comparability with GPP.EC when trees were
698 leafless in the rainy season. In both cases, GPP accounted only for crops (groundnut and cowpea)
699 and weeds, as trees were non-photosynthetic in the rainy season. Despite this standardisation,
700 GPP.EC values (-5.1 ± 3.6 g C-CO₂ m⁻² d⁻¹) were significantly higher than GPPch.stand values (-4.2
701 ± 4.3 g C-CO₂ m⁻² d⁻¹). However, the divergence did not occur on the peak of GPP (which was very
702 similar in both methods), but from the onset of groundnut senescence, when weeds became the
703 dominant photosynthetic contributors. Thus, during the groundnut growth season, with leafless
704 *F. albida* trees and almost no weeds, GPP measurements from EC and chambers generate closely
705 comparable results. Therefore, this provides an initial form of cross-validation between the two
706 methods. It is important to note that the EC method integrates CO₂ fluxes over a larger spatial
707 scale, encompassing all ecosystem components (Baldocchi, 2003), while the chamber method
708 captures fluxes on a smaller scale (i.e., at the 0.25 m² scale). This scale disparity can introduce
709 uncertainties when upscaling chamber-based fluxes to the field, as vegetation composition within
710 chambers does not represent the EC footprint's average vegetation. This makes upscaling
711 chamber-based measurements challenging. Nevertheless, the standardisation we applied on
712 chamber photosynthesis by LAI has been relatively successful.

713 During the dry season, Reco.EC included respiration from below- and above-ground tree parts
714 (with leaves) and bare soil, while Rch.stand measured only below-ground tree and bare soil



715 respiration. Consequently, the difference between Reco.EC and Rch.stand was solely attributable
716 to above-ground tree respiration (Ra tree above-ground). In terms of GPP, chamber
717 measurements were nil, whereas GPP.EC reflected only GPP trees.

718 The transition period, characterised by groundnut senescence, tree leaf regrowth, and weed
719 proliferation, introduced further complexity, amplifying method-specific discrepancies. Rch.stand
720 measurements facilitated the estimation of tree contribution to Reco.EC (Ra tree) and the
721 verification of the consistency for EC results in terms of carbon use efficiency (CUE), estimated
722 here at 0.48. This value indicates that nearly 50% of the carbon captured by trees is allocated to
723 biomass. The CUE estimate here is well comparable to the global average across diverse
724 ecosystems, climates, and management practices (0.49 ± 0.14) (Tang et al., 2019). Similar CUE
725 values have been reported for semi-arid grasslands (0.46 ± 0.10), but our value is notably lower
726 than those documented for wetlands (0.61 ± 0.13) (Tang et al., 2019). Overall, these findings
727 reinforce the plausibility of our assumptions regarding the compartment's contributions to
728 Reco.EC and Rch.stand, thereby providing a second cross-validation of the EC-Ch comparison.
729 However, despite a frequently assumed CUE of 0.5 in models, global estimates span a broad range
730 (0.20 to 0.82), depending on ecosystem type and management practices (DeLucia et al., 2007;
731 Tang et al., 2019). This underscores the importance of refining carbon flux models to better
732 represent the biophysical processes governing CO₂ exchange in semi-arid agroforestry systems.
733 The combined use of EC and chamber methodologies offers a comprehensive perspective on
734 ecosystem-scale CO₂ flux dynamics, advancing understanding of carbon cycling in these
735 environments.

736 4.6. Net carbon exchange budget

737 The annual net carbon (C) exchange budget was quantified at -1.4 ± 0.02 Mg C-CO₂ ha⁻¹ with the
738 chamber method and -1.8 ± 0.01 Mg C-CO₂ ha⁻¹ by the Eddy Covariance (EC), indicating that the
739 studied agro-silvo-pastoral system functions as a net carbon sink. These findings corroborate the
740 system's potential role in mitigating greenhouse gas emissions, consistent with previous
741 observations in semi-arid ecosystems (Rahimi et al., 2021; Tagesson et al., 2015; Agbohessou et
742 al., 2023, Wieckowski et al., 2024).

743 The estimated net C exchange budget is close to the reported mean for Sahelian ecosystems (-1.6
744 ± 0.5 Mg C-CO₂ ha⁻¹; Tagesson et al., 2016). The EC-based net C exchange budget (-1.8 ± 0.01 Mg
745 C-CO₂ ha⁻¹) is also similar to the value of -1.9 ± 0.4 Mg C-CO₂ ha⁻¹ reported for semi-arid savannas
746 of northeastern Benin, despite higher annual rainfall (1495 mm; Ago et al., 2016b). Furthermore,
747 our EC estimate is close to the average net C exchange reported for West African terrestrial
748 ecosystems (-2.0 ± 1.5 Mg C-CO₂ ha⁻¹; Ago et al., 2016a).



749 However, estimates from Tagesson et al. (2015) (-2.7 ± 0.07 Mg C-CO₂ ha⁻¹) for a semi-arid
750 savannah in Dahra, Senegal, located between the 300 mm and 400 mm isohyets, were
751 comparatively higher. This is potentially attributable to specific characteristics of that specific
752 savannah site, such as herbaceous vegetation cover during the rainy season, the presence of
753 evergreen trees, and land management practices linked to pastoral livestock activities (Tagesson
754 et al., 2016).

755 The net C exchange estimates presented in this study are, in fact, apparent fluxes, given that they
756 exclude organic matter (OM) imports and, more critically, exports, introducing uncertainties.
757 Notably, the export of crop residues and direct inputs from animal excreta —particularly
758 significant in ‘bush fields’ during the dry season — were not accounted for. In our case of ‘bush
759 field’, crop residues are exported to feed livestock, while livestock faeces are collected for use as
760 fuel or manure in ‘home fields’. Such practices may lead to a significant soil organic carbon stocks
761 depletion (Malou et al, 2021), potentially diminishing the net C budget (-1.4 ± 0.02 Mg C-CO₂ ha⁻¹)
762 over time and shifting the system closer to carbon neutrality (Assouma et al., 2019).

763 These results should be contextualized within the broader framework of climate change and semi-
764 arid ecosystem management. Although agro-silvo-pastoral systems can function as annual carbon
765 sinks, they remain highly sensitive to interannual rainfall variability and escalating anthropogenic
766 pressures. Sustainable management practices, particularly regarding crop residue exports, are
767 essential for maintaining soil mineral fertility and preserving the system’s capacity to act as a
768 carbon sink, thereby contributing to climate change mitigation.

769 4.7. Limitations of the study

770 This study benefited from the inverse phenology of *F. albida*, allowing for direct comparison
771 between chamber-based GPP (GPPch.stand) and ecosystem-level GPP (GPP.EC) during the
772 leafless period of the trees. However, the system’s spatial heterogeneity —common in
773 agroforestry— posed challenges for accurately partitioning CO₂ fluxes among trees, crops, and
774 soil. A key limitation was the development of weeds during the late rainy season, which
775 complicated the attribution of fluxes, particularly during the transitional period. Additionally,
776 while GPPch was successfully standardised by LAI for upscaling, this was not feasible for
777 respiration. Respiration integrates both autotrophic and heterotrophic components, which
778 respond to different drivers and are not directly linked to LAI, limiting the precision of upscaled
779 Rch.

780 Future improvements should aim to separately quantify respiration sources —tree roots, crops,
781 and microbial (heterotrophic) respiration— and account explicitly for the weed layer, to refine
782 flux partitioning in such complex agroforestry systems.



783 **Conclusion**

784 This study demonstrates the successful application of automated static chambers to quantify CO₂
785 fluxes in a Sahelian agroforestry system dominated by *F. albida*. The continuous, high-frequency
786 measurements captured key seasonal dynamics and short-lived events (e.g., Birch effect),
787 providing a more accurate assessment of carbon exchange than traditional intermittent sampling.
788 By integrating crop and soil components and applying dynamic partitioning models, the study
789 quantified both respiration and photosynthesis fluxes at fine temporal resolution. The results
790 revealed a clear 'fertile island' effect under tree canopies, with higher respiration and
791 photosynthetic activity, and highlighted the significant contribution of *F. albida* trees to annual
792 carbon uptake.
793 The consistency between chamber- and eddy covariance-based estimates reinforces the
794 robustness of the methodology. Overall, this work underscores the role of *F. albida*-based
795 agroforestry systems as effective carbon sinks in semi-arid environments, offering valuable
796 insights for carbon accounting and sustainable land management in the Sahel.



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810 **Author contribution: CRediT**

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813 apparatus and methodology, writing, review and editing. **Lydie Chapuis-Lardy:** Designing
814 methodology, writing, review and editing. **Yélognissè Agbohessou:** Processing data,
815 review and editing. **Fred Bouvery:** Designing chambers and connection to the instrument,
816 review and editing. **Maxime Duthoit:** Designing experimental set and methodology,
817 review and editing. **Aleksander Wieckowski:** Review and editing. **Mohamed Habibou**
818 **Assouma:** Review and editing. **Espoir Gaglo:** Processing data, review and editing. **Claire**
819 **Delon:** review and editing. **Torbern Tagesson:** Designing methodology, review, and
820 editing. **Bienvenu Sambou:** Review and editing. **Dominique Serça:** Designing methodology,
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