

1 **Drivers and vertical CO₂ flux balances budgets in a Sahelian *Faidherbia albida* agro-silvo-**
2 **pastoral parkland: Insights from continuous high-frequency soil chamber measurements**
3 **and Eddy Covariance.**

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32 **Highlights:**

- 33 • Long-term high frequency CO₂ flux measurements using automated static
34 chambers in a Sahelian *F. albida* parkland.
- 35 • Empirical gap-filling and flux partitioning methods validated against Eddy
36 Covariance GPP.
- 37 • Fluxes peaked during the rainy season ~~in~~ both at a distance from trees in full sun
38 (FS) and under tree canopies (Sh), driven mainly by soil moisture and leaf area.
- 39 • *F. albida* trees enhance CO₂ fluxes under canopies ("fertile island" effect) and
40 account for ~23% ~~~50%~~ of annual ecosystem GPP.

41 **ABSTRACT:**

42 Agroforestry systems — combining trees with crops and/or livestock — are increasingly
43 promoted as sustainable and climate-resilient land-use strategies. Despite their widespread
44 presence in the Sahel, experimental data on their potential as carbon sinks are scarce. This study
45 presents a full-year, high-frequency dataset of CO₂ fluxes in a Sahelian agro-silvo-pastoral
46 parkland dominated by *F. Faidherbia albida*, located in Senegal's groundnut basin. CO₂ fluxes were
47 continuously measured using automated static chambers, allowing the quantification of soil and
48 crop respiration (R_{ch}), gross primary production (GPP_{ch}), and net carbon exchange (FCO₂ch)
49 under both full sun and shaded (under tree canopies) environments.

50 Seasonal patterns of CO₂ fluxes were similar in both environments, with peaks during the rainy
51 season. R_{ch} and GPP_{ch} were significantly higher under tree canopies, indicating a 'fertile island'
52 effect. CO₂ flux variability was primarily driven by soil moisture and leaf area index. Chamber-
53 based GPP estimates closely matched those from Eddy Covariance measurements. On an annual
54 scale, *F. albida* trees contributed approximately ~~23%~~ ~~50%~~ of total ecosystem GPP, with a carbon
55 use efficiency of 0.48. Net annual ~~vertical~~ CO₂ exchange was estimated at -1.4 ± 0.4602 and -1.8
56 ± 0.1704 Mg C-CO₂ ha⁻¹ using chamber and Eddy Covariance methods, respectively. These
57 findings underscore the role of *F. albida*-based agroforestry systems as effective carbon sinks in
58 Sahelian landscapes, supporting their potential contribution to climate change mitigation.

59 **Keywords:** Sahelian agro-silvo-pastoral systems, CO₂ fluxes, automated static chambers, Eddy
60 Covariance, 'fertile island effect' ~~of trees~~, carbon ~~balances~~ ~~budgets~~.

61 **1. Introduction**

62 Plant photosynthesis and respiration —both autotrophic (plant) and heterotrophic (microbial)—
63 are fundamental processes driving carbon dioxide (CO₂) fluxes in terrestrial ecosystems
64 (Lambers et al., 2008; Raich et al., 2014; Reichle, 2020). Accurate quantification of these processes
65 is critical for assessing ecosystem carbon (C) sink potential (Baldocchi, 2020), particularly for
66 informing climate-smart land management strategies.

67 To capture these processes at the ecosystem scale, the Eddy Covariance (EC) technique has
68 emerged as a transformative method, enabling continuous and high-frequency CO₂ flux
69 measurements (Baldocchi, 2003, 2008). The EC technique quantifies CO₂ exchanges between
70 ecosystems and the atmosphere by correlating fluctuations in vertical wind velocity with
71 simultaneous variations in CO₂ concentrations, providing a direct and non-invasive estimate of
72 CO₂ fluxes (Baldocchi, 2003). Extensive EC networks in Europe (Stojanović et al., 2024), Asia (Yu
73 et al., 2011), and the Americas (Chu et al., 2021) have significantly advanced our understanding
74 of the global C cycle. In contrast, sub-Saharan Africa remains critically underrepresented
75 (Bombelli et al., 2009; Houghton & Hackler, 2006; Williams et al., 2007). Although some studies
76 have used EC (Ardö et al., 2008; Brümmer et al., 2008; Merbold et al., 2009; Tagesson et al., 2016),
77 static chambers (Assouma et al., 2017; Owusu et al., 2024; Rosenstock et al., 2016; Wachiye et al.,
78 2020), or modeling approaches (Agbohessou et al., 2023, 2024; Delon et al., 2019; Rahimi et al.,
79 2021), they remain sparse and methodologically heterogeneous, limiting comparability and
80 regional C budget integration.

81 Among these underrepresented landscapes, agroforestry systems in the Sahel— particularly
82 agro-silvo-pastoral systems (ASPS) that combine trees, crops, and livestock— are increasingly
83 promoted for sustainable land management and climate resilience (Cardinael et al., 2021; Gupta
84 et al., 2023; Mbow et al., 2014; Stetter & Sauer, 2024). However, the structural and functional
85 heterogeneity of these systems poses significant challenges for accurately quantifying and
86 upscaling C fluxes. *Faidherbia albida*, a keystone agroforestry tree species in these ASPs (Leroux
87 et al., 2022; Lu et al., 2022), is of particular interest due to its reverse phenology, capacity to
88 enhance soil fertility and crop yields (Bayala et al., 2020; Roupsard et al., 2020; Sileshi et al., 2016;
89 2020). Yet, its functional role in modulating both the magnitude and seasonal dynamics of CO₂
90 fluxes remains poorly understood.

91 Addressing this knowledge gap requires integrated approaches capable of capturing both
92 aggregate and component-specific CO₂ fluxes. While EC remains the gold standard method for CO₂
93 flux measurements at the landscape scale (Baldocchi, 2003), it captures net ecosystem exchange
94 (NEE) as an aggregate signal, without separating the contributions from individual compartments
95 such as soil, crops, and trees. This limits its utility for disentangling processes and attributing
96 sources in heterogeneous systems like ASPs. Automatic static chambers provide a valuable

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97 complement to EC, as they enable continuous, high-frequency measurements at finer scales and
98 at the level of specific ecosystem components. This approach facilitates component-specific
99 quantification of CO₂ fluxes, particularly from soil and crop compartments (Luo & Zhou, 2006;
100 Denmead, 2008; Zaman et al., 2021). When combined with EC, this dual-method approach
101 strengthens source attribution and improves the partitioning upscaling of fluxes across complex
102 agroforestry landscapes.

103 This study presents one of the first integrated quantification of CO₂ fluxes in a Sahelian ASPS
104 dominated by *F. albida*, combining EC and automatic static chambers.

105 Specifically, we the study aims to (1) conduct year-round, high-frequency *in situ* CO₂ flux
106 measurements from soil and crops using automated static chambers; (2) partition the net CO₂
107 fluxes (FCO₂ch) into respiration (Rch) and photosynthesis (GPPch); (3) investigate the
108 environmental drivers of fluxes and the spatial variability linked to tree presence; and (4)
109 compare chamber-based flux estimates with ecosystem-scale measurements derived from the EC
110 method.

111 Based on these objectives, we hypothesize that (1) Rch and GPPch are higher under the canopy of
112 *F. albida* than in full sun, (2) soil moisture is the main environmental factor directly controlling
113 both Rch and GPPch, (3) when extrapolated to the field scale, the chamber-based method provides
114 seasonal dynamics of respiration and photosynthesis fluxes comparable to those derived from EC
115 technique.

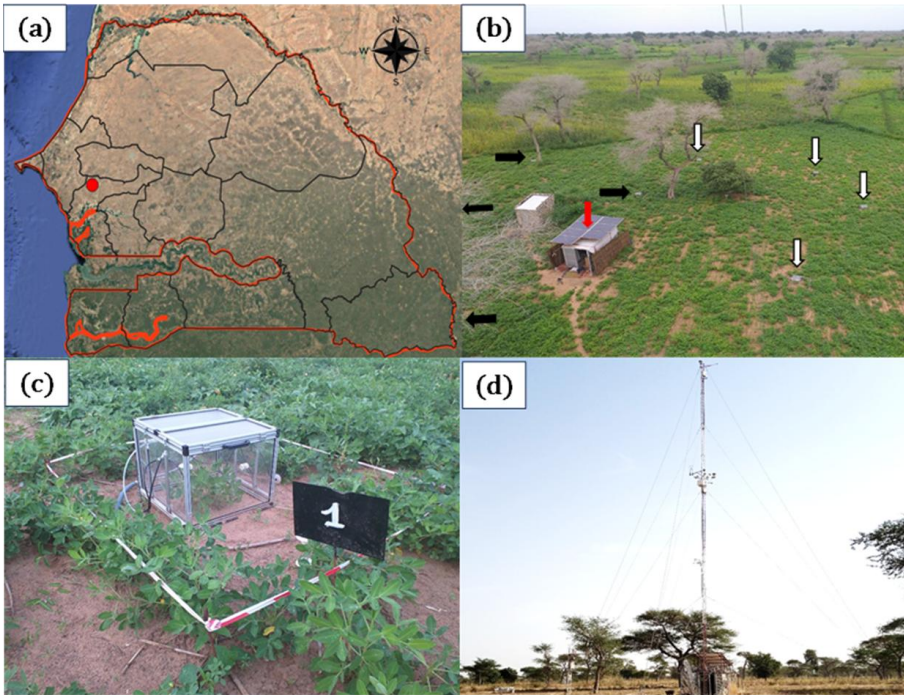
116 **2. Materials and methods**

117 *2.1. Site description*

118 The study was conducted in the agroforestry parkland of Sob village (Niakhar municipality, Fatick
119 region), located in the groundnut basin of Senegal, within the Sahelo-Sudanian climatic zone of
120 West Africa (Fig. 1). The climate is characterized by a long dry season (8–9 months) with high
121 temperatures and strong diurnal variations, and a short rainy season from late June to mid-
122 October (Delaunay et al., 2018).

123 Soils are locally known as "*Dior*" and classified as Arenosols (IUSS Working Group WRB, 2022).
124 The topsoil has low organic matter (<1%) and phosphorus (<3 mg kg⁻¹), a sandy texture (>85%
125 sand), and an acidic pH (Malou et al., 2021; Siegwart et al., 2022). Rainfed agriculture
126 predominates. The main cropping system includes pearl millet (*Pennisetum glaucum L.*) and
127 groundnut (*Arachis hypogaea L.*) in biennial rotation, with occasional intercropping of cowpea
128 (*Vigna unguiculata L.*).

129 The site hosts the 'Faidherbia Flux' station (14°29'44.916"N; 16°27'12.851"W; FLUXNET ID: SN-
130 Nkr), a long-term research platform for monitoring ecosystem services in agroforestry systems.
131 It is dominated by *F. albida*, a nitrogen-fixing, reverse-phenology tree with deep roots accessing
132 groundwater (Roupsard et al., 1999). The tree density is ~13 trees ha⁻¹, with canopies covering
133 ~10% of the soil surface (Roupsard et al., 2020). The EC tower is installed at 20 m height,
134 approximately 12.5 m above the canopy. The study field is a typical 'bush field', characterized by
135 low soil fertility, no mineral fertilization, and off-site export of crop residues and manure (Malou
136 et al., 2021).



137 Fig. 1: Study area.

138 (a) geographical location of Sob, Groundnut basin, Senegal (Map data © Google Earth, 2025), (b) overview
 139 (image from the Eddy Covariance tower located in the same bush-field) of the *Faidherbia albida* parkland
 140 during the rainy season, depicting groundnut crops with bare soil in the inter-row, *F. albida* trees
 141 (defoliated during the rainy season, average height = 13m) and location of the chambers under the Shade
 142 of trees (horizontal black arrows; N=4) and in Full sun (vertical white arrows; N=4); The shelter (red
 143 arrow) with solar panels is to fit the analyser, automation and batteries (c) automatic chamber enclosing a
 144 groundnut plant (during the rainy season) or bare soil (during the dry season), (d) Eddy Covariance (EC)
 145 tower (measurement height = 20 m) during the dry season.

146 2.2. *Experimental setup*

147 2.2.1. *CO₂ flux measurements in automatic chambers*

148 Continuous net CO₂ fluxes (FCO₂ch) from soil and groundnut plants were measured over a full
149 phenological year (June 17, 2021 – June 17, 2022) using eight automated static chambers
150 (50×50×50 cm), each enclosing one groundnut plant. Four chambers were installed in full sun
151 (FS), at least 20 m from trees, and four under *F. albida* canopy shade (Sh). The chambers were
152 transparent, custom-built (Duthoit et al., 2020), and installed on metal bases embedded 10 cm
153 into the soil one month prior to measurements.

154 During the rainy season (June–November), groundnut coexisted briefly with spontaneous weeds
155 until weeding (mid-July), after which chambers contained only groundnut. Post-harvest (early
156 November), chambers remained bare while surrounding plots experienced weed regrowth.

157 CO₂ concentrations were measured at 1 Hz using a Picarro G2508 gas analyser (Picarro Inc., Santa
158 Clara, CA, USA) (Fleck et al., 2013; Reum et al., 2019; Valujeva et al., 2022). A fully automated
159 system was built for sequential half-hour flux measurements ~~(alternating FS and Sh chambers)~~
160 [\(Table S2\)](#). Measurement duration was 15 min per chamber in the dry season, reduced to 5 min
161 during the rainy season to limit condensation effects.

162 2.2.2. *CO₂ flux measurements by Eddy Covariance*

163 The EC system (Li-COR SMARTFLUX®, including a Gill MasterPro 3D sonic anemometer and a LI-
164 7500 RS open path CO₂ and H₂O gas analyser) was mounted at a height of 20 m on a 30m mast,
165 above *F. albida*. It continuously monitored net CO₂ exchange from the ecosystem. Raw data were
166 collected at 20 Hz frequency and post-processed from binary files using the advanced mode of the
167 EddyPro® v7.0, with standard corrections and procedures: sonic tilt correction (double rotation),
168 block averaging, covariance maximisation for time lag, and WPL correction (Webb et al., 1980).
169 Quality control followed Foken et al. (2004) and Vickers & Mahrt (1997); random uncertainty was
170 estimated per Finkelstein & Sims (2001). Spectral corrections were applied according to
171 Moncrieff et al. (1997, 2004). Footprints were computed according to Kormann and Meixner
172 (2001), using the FREddyPro R package (Xenakis, 2016), ~~indicating indicated~~ a ~1 ha source area
173 covering the entire field. Gap-filling and flux partitioning were conducted using ReddyProc
174 (Wutzler et al., 2018), applying the daytime partitioning approach of Lasslop et al. (2010).

175 2.2.3. *Ancillary measurements*

176 Environmental and vegetation variables were monitored continuously throughout the study.
177 Global radiation (R_g) was estimated from photosynthetically active radiation (PAR) using a Skye
178 sensor (averaged over 30-min intervals). [The normalised difference vegetation index \(NDVI\)](#) of

179 crops under full sun was recorded semi-hourly by a calibrated downward-facing sensor installed
180 at 20 m height (Pontauiller et al., 2003), processed following Soudani et al. (2012), and used to
181 estimate [the leaf area index \(LAI\)](#) time series for groundnut, weeds, and cowpea based on end-of-
182 season field LAI measurements in six 15 m² plots (as in Roupsard et al., 2020).

183 Rainfall was recorded by an automatic weather station (CR1000 with TE525MM rain gauge,
184 Campbell Scientific), and soil volumetric water content (VWC) and temperature (T_{soil}, at 6 cm
185 depth) were monitored using TOMST® TMS-4 sensors, benchmarked prior to field deployment
186 inside and outside the chambers (Wild et al., 2019). Air temperature (T_{air}) was recorded inside
187 each chamber at 15 cm above ground, all at 5-min intervals. These measurements contribute to
188 the SoilTemp global database (Lembrechts et al., 2020, 2022).

189 Groundnut development was tracked weekly by counting leaves in each chamber. Total
190 groundnut LAI (LAI_{ch}) was then derived from average single-leaf area and chamber surface.

191 A detailed description of the data used in this study is provided in Supplement [S1](#) (Table [S1.1](#)).

192 2.3. Data processing

193 2.3.1. Flux calculation

194 Net CO₂ fluxes (FCO_{2ch}, in μmol CO₂ m⁻² s⁻¹) from the chambers were calculated from the linear
195 change in CO₂ concentration over time (ΔC/Δt; [Fig. S1 and Fig. S2](#)) using the Eq.1.

$$196 \text{FCO}_{2\text{ch}} = \left(\frac{P}{RT_k}\right) \left(\frac{V}{A}\right) \left(\frac{\Delta C}{\Delta t}\right) \quad (\text{Eq. 1})$$

197 where P is atmospheric pressure (101 325 N m⁻²), R is the ideal gas constant (8.31 N m mol⁻¹ K⁻¹),
198 T_k is air temperature inside the chamber in Kelvin, V (0.125 m³) is the total system volume
199 (chamber, tubing, analyser cavity, pump, and water trap), and A (0.25 m²) is the chamber
200 footprint. The slope ΔC/Δt was obtained via linear regression (Duthoit et al., 2020).

201 Mean FCO_{2ch} values were computed separately for the four replicate chambers in full sun (FS)
202 and under *F. albida* shade (Sh). By convention, negative values indicate net CO₂ uptake
203 (photosynthesis), and positive values indicate net CO₂ release (respiration).

204 2.3.2. Quality control of chamber-based CO₂ flux measurements

205 The quality of chamber-based CO₂ flux measurements was assessed using [a threshold of the](#)
206 [coefficient of determination \(R² ≥ 0.8\)](#) of the linear increase in CO₂ concentration during chamber
207 closure. The minimum detectable flux (MDF) was then calculated following Nickerson (2016)
208 (Eq.2). The MDF defines the flux detection threshold, below which data are considered unreliable
209 due to instrument sensitivity and sampling constraints (Zaman et al., 2021). In this study, the MDF
210 was ±0.0004 μmol CO₂ m⁻² s⁻¹.

211
$$\mathbf{MDF} = \left(\frac{A_a}{t_c(\sqrt{t_c/p_s})} \right) \left(\frac{VP}{ART} \right) \quad \mathbf{(Eq. 2)}$$

212 where A_a is the analytical precision of the Picarro analyser (0.6 ppm; Picarro Inc., 2015), t_c the
213 closure time (s), p_s the sampling frequency (1 Hz), V the chamber volume, P the atmospheric
214 pressure (101 325 N m⁻²), A the chamber footprint, R the gas constant (8.3 N m mol⁻¹·K⁻¹), and T
215 the air temperature in Kelvin.

216 Following this quality control, fluxes were partitioned (Section 2.3.3) and gap-filled (Section
217 2.3.4).

218 *2.3.3. Partitioning of chamber-based CO₂ fluxes*

219 The net CO₂ fluxes (FCO₂ch), averaged from four chambers per environment (FS and Sh), were
220 partitioned into two components according to Eq. 3 (Reichstein et al., 2005).

221
$$\mathbf{FCO_2ch} = \mathbf{Rch} + \mathbf{GPPch} \quad \mathbf{(Eq. 3)}$$

222 Rch includes heterotrophic respiration (Rh) from soil and other autotrophic respiration (Ra) from
223 groundnut plants and roots of *F. albida* (Ra Groundnut + Ra tree below-ground). Rch is always
224 positive (Rch > 0). GPPch (Gross Primary Productivity) represents the photosynthetic CO₂ uptake
225 by the groundnut plants and is negative during the day (GPPch < 0), and zero at night, when
226 FCO₂ch = Rch.

227 Half-hourly FCO₂ch fluxes were partitioned as follows: (1) an Arrhenius-type function (Lloyd &
228 Taylor, 1994) was fitted between nocturnal Rch and T_{soil} during nighttime periods, for each 5-days
229 throughout the time series (Eq. 4). This empirical formulation is based on several key
230 assumptions. First, the relationship between nocturnal respiration and soil temperature is
231 assumed to follow an exponential response, reflecting the temperature sensitivity of respiration
232 processes. Second, the model assumes temporal stability of the respiration–temperature
233 relationship between night and day, allowing diurnal respiration to be extrapolated from fitted
234 parameters in Eq.4 and daytime T_{soil}. Third, we assumed that no abrupt changes in substrate
235 availability or soil moisture occur between day and night — conditions that could otherwise
236 disrupt the temperature–respiration relationship. ~~Third, it is assumed that no abrupt changes in~~
237 ~~substrate availability or soil moisture occur between night and day — conditions that could~~
238 ~~otherwise decouple respiration rates from temperature.~~ These assumptions are widely applied in
239 CO₂ flux partitioning approaches (Reichstein et al., 2005; Lasslop et al., 2010). (2) Diurnal Rch
240 was estimated by applying the Lloyd & Taylor function, previously calibrated on nocturnal data,
241 to the corresponding daytime T_{soil} measurements for each 5-day interval. (3) GPPch was
242 subsequently derived as the residual component of the net CO₂ flux during the day, according to:

243 **nocturnal Rch** = $R_{\text{ref}} \cdot \exp \left[E_0 \left(\frac{1}{T_{\text{ref}} - T_0} - \frac{1}{T_{\text{soil}} - T_0} \right) \right]$ (Eq. 4)

244 where R_{ref} ($\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$) is a fitted parameter representing the base respiration at the
 245 reference temperature [T_{ref} (K), (set at 288.15 K)]. E_0 (K) is the temperature sensitivity (set at
 246 250 K), T_{soil} (K) the soil temperature (K), and T_0 (K) is kept constant at 231.13 K, according to
 247 Lloyd & Taylor (1994).

248 **GPPch** = **diurnal FCO₂ch** – **diurnal Rch** (Eq. 5)

249 where diurnal FCO₂ch and diurnal Rch represent the daytime net CO₂ fluxes and respiration in
 250 $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$, respectively.

251 2.3.4. Gap-filling procedure

252 Missing Rch data were gap-filled using the model derived from Eq. 4 (Lloyd & Taylor, 1994). Prior
 253 to gap-filling GPPch, raw data were standardised by LAI to reduce variability between chambers
 254 due to differences in leaf surface area (Eq. 6). A light-response model was then fitted to the
 255 standardised GPPch data, every 5-day period, to gap-fill missing values. The model is based on a
 256 rectangular hyperbolic function that describes the relationship between photosynthetic CO₂
 257 uptake and incoming global radiation (Rg) (Eq. 7). It corresponds to a Michaelis–Menten-type
 258 light-response curve, commonly used in ecosystem carbon exchange studies (Falge et al., 2001;
 259 Lasslop et al., 2010).

260 **GPPch.stand** = $\frac{\text{GPPch}}{\text{LAIch}} * \text{LAI.field}$ (Eq. 6)

261 where GPPch.stand ($\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$) is the standardised GPPch. LAIch and LAI.field (m^2 leaves
 262 m^{-2} soil) represent the groundnut LAI inside the chambers and the groundnut + weeds + cowpea
 263 LAI for the whole field, respectively.

264 **GPP** = $\frac{\alpha\beta Rg}{\alpha Rg + \beta}$ (Eq. 7)

265 where α ($\mu\text{mol CO}_2 \text{ J}^{-1}$) represents the light use efficiency of the groundnut plants inside the
 266 chambers, and refers to the initial slope of the light-response curve, β ($\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$) is the
 267 maximum CO₂ uptake rate by the groundnut plants at light saturation, and Rg the global radiation
 268 (W m^{-2}).

269 2.3.5. Comparing chamber-based (Ch) and Eddy Covariance (EC) methods

270 Chamber measurements were upscaled to field-level CO₂ fluxes and compared with EC-derived
 271 fluxes. Before comparison, a correction was applied (Eq. 6) to account for differences in LAI
 272 between chambers (LAIch) and the field (LAI.field), due to the presence of cowpea and weeds in
 273 the field but not in the weeded chambers.

274 Upscaling considered tree cover, with FS and Sh chamber fluxes weighted at 90% and 10%,
 275 respectively. Rch.stand and GPPch.stand, representing chamber-based respiration and

276 photosynthesis at field scale. These fluxes were compared, on a half-hourly basis, to EC-derived
277 Reco.EC and GPP.EC (S3, Table S43.1). The November–December transition period was excluded
278 due to weed-driven uncertainties after groundnut harvest.

279 During the rainy season (*F. albida* leafless), GPP.EC represented ground vegetation (groundnut,
280 cowpea, weeds), while Reco.EC included autotrophic respiration from all vegetation (including
281 trees), and heterotrophic respiration (Reco.EC = Ra tree below-ground + Ra tree above-ground
282 + Ra groundnut + Ra cowpea + Ra weeds + Rh). Rch.stand could not be fully upscaled to the field
283 due to uncertainty in its partitioning between Ra and Rh. Rch.stand accounted only for Ra tree
284 below-ground, Ra groundnut, and Rh.

285 In the dry season (leafy trees, bare soil), GPP.EC reflected tree photosynthesis only (GPP tree),
286 while GPPch.stand was nil. Reco.EC included Ra tree (above- and below-ground) and Rh.
287 Rch.stand, measured on bare soil, represented only Ra tree below-ground + Rh.

288 2.3.6. Contribution of trees to full ecosystem respiration and photosynthesis

289 During the dry season, when the trees (*F. albida*) maintained their foliage, a comparison between
290 chamber and EC measurements allowed for the estimation of the contribution of the above-
291 ground tree compartments to total ecosystem respiration (S3, Table S43.1). Based on this
292 estimate, total tree respiration (Ra tree) was then calculated under the assumption that the tree
293 root systems (Ra tree below-ground) represent 1/3 of the above-ground biomass (Jackson et al.
294 1996).

295 Given the GPP measured during the dry season was equivalent to GPP of trees (GPP trees) from
296 EC measurements, the carbon use efficiency of the trees (CUE tree) was then calculated (S3, Table
297 S43.1). The resulting CUE value was assessed to determine whether it approximated the typical
298 value of 0.5, which is often used as a default in ecosystem models (Zhou et al., 2019; 2020).

299 2.3.7. Net annual *vertical C balance budget* at the ASPs scale

300 The ~~net annual-C balance budget~~ of CO₂ fluxes ~~in a yearly basis~~ was estimated for chambers and
301 EC measurements in Mg C-CO₂ ha⁻¹. The chambers CO₂ ~~flux balances fluxes budgets~~ were
302 obtained by calculating the annual sum of the net CO₂ flux measurements and then weighting with
303 the tree cover rate (10% for the Sh, 90% for the FS). These annual ~~budgets balances~~ for the field
304 are considered apparent ~~representing vertical CO₂ exchanges only~~, as they do not account for the
305 biomass exported from the field after the harvest, the decomposition of which therefore escaped
306 both the chambers and the EC. Additionally, the inputs and the outputs of fecal matter resulting
307 from livestock wandering during the dry season were not quantified and are therefore neglected.
308 The objective here is to compare two approaches at different scales using ~~apparent-vertical~~ net C

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309 ~~budgets~~balances, rather than to provide an absolute C budget which would also include horizontal
310 transfers of carbon.

311 2.4. Statistical analyses

312 Statistical analyses were performed using the R software (R. Core Team, 2023). To compare the
313 mean values of climatic parameters between the FS and Sh situations, a non-parametric Mann-
314 Whitney test was used when both the normality (shapiro.test) and the homogeneity of the
315 variance (Levene Test, R package 'Car'; Fox et al., 2023) were not confirmed. This approach was
316 similarly applied to compare the seasonal dynamics of CO₂ fluxes between FS and Sh, as well as
317 between the chamber-based and Eddy Covariance (EC) methods. Means and standard deviations
318 were computed using the 'skim' function from the R package 'skimr' (Waring et al., 2022).

319 Respiration (Rch) (Eq. 4) and GPP (GPPch) models (Eq. 7) were fitted using non-linear least
320 squares regression, implemented in the library in R 'nls.multstart' (Padfield et al., 2025). For the
321 GPPch model, parameters α and β with non-significant p-values were removed, and then the
322 remaining values were interpolated and smoothed using a 'spline' function from the 'zoo' library
323 in R (Zeileis et al., 2024). Ordinary least-square linear regressions were fitted between the
324 measured and the modeled values derived from. Model performance of Eq. 4 and Eq. 7 was
325 evaluated by fitting ordinary least-square linear regressions between the measured and the
326 modeled values using R², root mean square error (RMSE), and the bias metrics. Given that the
327 primary objective of these equations was to accurately reproduce the seasonal dynamics of the
328 CO₂ fluxes to fill gaps in data, particular emphasis was placed on R², with a higher value reflecting
329 a better fit of the model to the measurements.

330 Correlation analysis was conducted between chamber CO₂ fluxes (FCO₂ch, Rch, GPPch) and soil
331 temperature (T_{soil}, °C), air temperature (T_{air}, °C), VWC, the leaf area index of groundnut plants in
332 the chambers (LAIch), and the fitted parameters for respiration — R_{ref} — and photosynthesis —
333 α and β . This analysis was performed using the 'cor.test' function from the 'stats' package in R
334 (Lüdtke et al., 2021), applying the Spearman method.

335 The threshold of the daily mean soil temperature (T_{soil}, °C) at which the cumulative daily
336 respiration (Rch, g C-CO₂ m⁻² d⁻¹) began to decline was determined using segmented regression
337 from the R package 'segmented' (Muggeo, 2003). The associated uncertainty (standard error) of
338 this estimate was evaluated through a bootstrap procedure.

339 The standard error of the total annual flux was estimated using the error propagation method.
340 This calculation considered the mean standard deviation of daily fluxes (g C-CO₂ d⁻¹) and the
341 effective number of measurement days (365). For each FS and Sh condition, the mean daily
342 standard deviation was multiplied by the square root of 365 to obtain the annual standard error.

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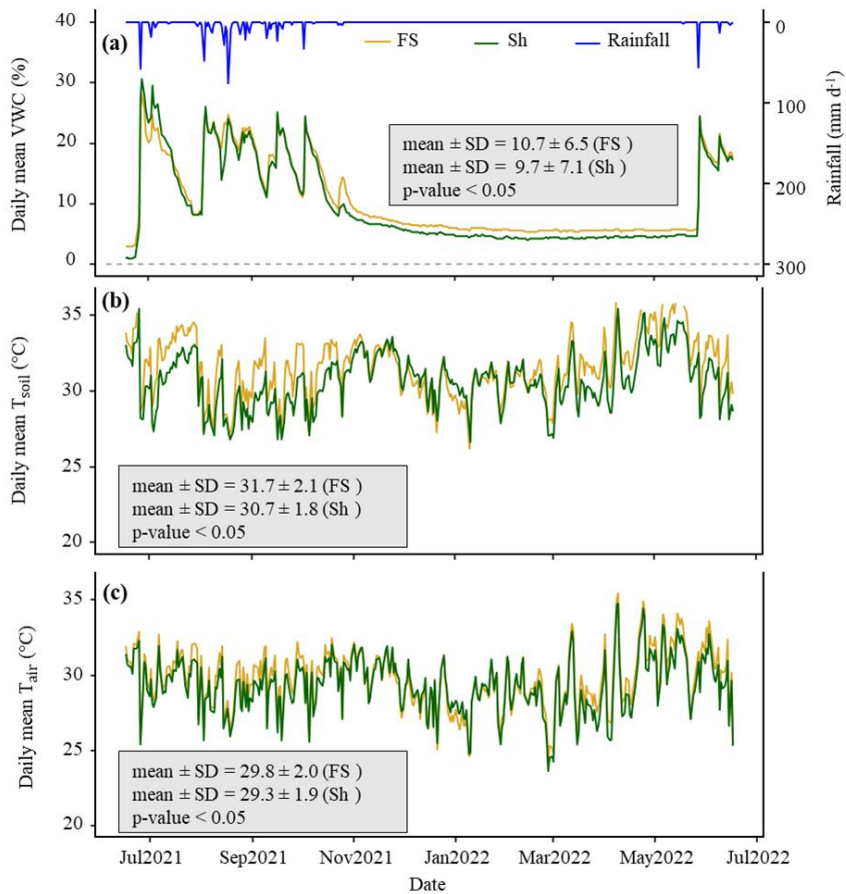
343 The resulting values were then weighted by 90% for FS and 10% for Sh to derive the overall
344 standard error of the annual flux sum, which was subsequently converted to Mg C-CO₂ ha⁻¹.

345 3. Results

346 3.1. Microclimatic conditions

347 During the experiment, the cumulative rainfall was 550 mm, which was representative of the
348 interannual average. Precipitations were lowest in July and highest between August and
349 September, a period that typically corresponds to the peak of the rainy season (Fig. 2a). Global
350 radiation ranged between 5.8 and 32.4 MJ m⁻² d⁻¹ (data not shown). The daily mean VWC in the
351 chambers showed significant variation, ranging from 1% at the end of the dry season to a
352 maximum of 30% during the rainy season (Fig. 2a). While VWC was similar during the rainy
353 season, it remained consistently higher in FS than in Sh throughout the dry season ($p < 0.05$),
354 which was unexpected. However, it should be noted that the last rain of October 2021 recharged
355 the FS chambers more effectively, likely due to foliage rainfall interception by *F. albida* which had
356 just put on leaves at that time, potentially explaining this discrepancy in VWC.

357 Within the chamber, the daily mean T_{soil} ranged from 26°C in April to 37.5°C at the end of the dry
358 season (Fig. 2b), while T_{air} varied between 23.7°C and 35.5°C (Fig. 2c). However, during
359 instantaneous daily peaks, T_{soil} could exceed 45°C in May (data not shown). As expected, both daily
360 mean T_{soil} and T_{air} were significantly higher in FS compared to Sh situations ($p < 0.05$), with T_{soil}
361 and T_{air} averaging respectively 1°C and 0.5°C lower under the tree canopy.



362 Fig. 2: One-year time series of daily average microclimatic parameters measured inside [the](#)
 363 chambers.

364 (a) volumetric soil water content (VWC) at a depth of 6 cm (%). (b) soil temperature (T_{soil}) at a depth of 6
 365 cm (°C), (c) air temperature (T_{air}) at a height of 15 cm (°C). The blue line depicts the daily rainfall (mm d⁻¹)
 366 throughout the year. FS: Full sun chambers; Sh: Shaded chambers. Mean and SD represent respectively the
 367 mean value and the standard deviation. The p-value indicates the probability associated with the statistical
 368 test, assessing the differences in means between FS and Sh with the significance level α set to 0.05.

369 3.2. Modeling the chamber-based total respiration (R_{ch}) and photosynthesis (GPP_{ch})

370 3.2.1. Dynamics of reference respiration, light use efficiency, and maximum CO_2 uptake rate at
371 light saturation (R_{ref} , α , and β)

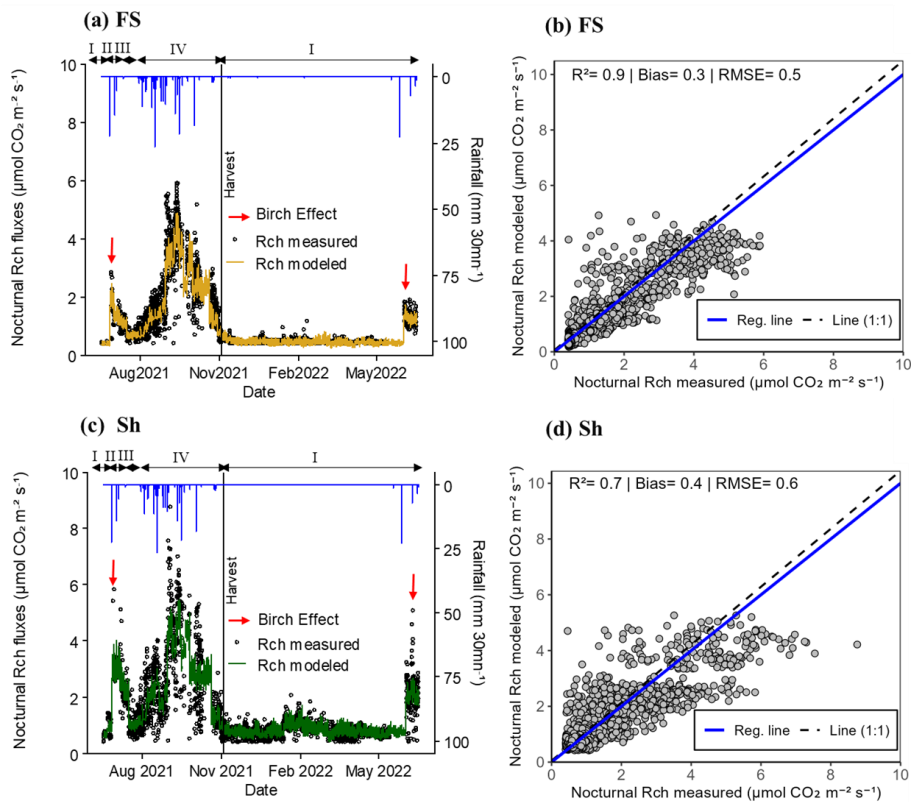
372 The reference respiration (R_{ref}) showed comparable seasonal dynamics both at a distance from
373 the trees (FS) and under the tree canopies (Sh) (S2, Fig. S42-2). In both situations, R_{ref} showed
374 strong variability during the rainy season, peaking in September 2021 at $2.4 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ for
375 FS and $2.9 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ for Sh (S2, Table S32-1). In contrast, during the dry season — from
376 November 3, 2021 (after harvest) until the onset of the following rainy season (June 2022) — R_{ref}
377 values dropped both for FS and Sh, averaging $0.3 \pm 0.5 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ for FS and $0.5 \pm 0.6 \mu\text{mol}$
378 $\text{CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ for Sh. This represents a reduction by a factor of 8 for FS and 6 for Sh compared to the
379 rainy season. The mean annual R_{ref} values were significantly higher under Sh than in FS, with value
380 approximately 1.5 times greater (S2, Table S32-1).

381 Regarding GPP in chambers, the light use efficiency (α) and the maximum CO_2 uptake by
382 groundnut plants in the chambers (β), also reached their maximum during the peak of the rainy
383 season (S2, Fig. S52-3, a and b). The maximum value of α reached $0.2 \mu\text{mol CO}_2 \text{ J}^{-1}$ in FS and 0.3
384 $\mu\text{mol CO}_2 \text{ J}^{-1}$ in Sh (S2, Table S32-4). Similarly, the maximum values of optimum CO_2 uptake rate
385 at light saturation (β) were $40.2 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ for FS and $42.8 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ for Sh (S2,
386 Table S32-4). In the dry season, when photosynthetic activity ceased in the chambers, both α and
387 β were assumed to be nil (S2, Fig. S52-3, a and b). On average, α and β were significantly higher in
388 Sh than in FS, by a factor of 1.7 and 1.2, respectively (S2, Table S32-4). We noted that the decline
389 in photosynthetic activity of the groundnut crop occurred earlier and rapidly at a distance from
390 the trees (FS), as reflected by the sharply observed recession of α and β in FS.

391 3.2.2. Dynamics of nocturnal respiration in chambers

392 The averaged nocturnal respiration (nocturnal R_{ch}) calculated from the measurements across
393 each treatment (FS and Sh), showed similar seasonal patterns (Fig. 3, a and c). Following the first
394 rains, R_{ch} values increased dramatically, with a nocturnal 'Birch effect' — a sudden pulse of CO_2
395 release following soil rewetting — observed to be more pronounced under Sh compared to FS,
396 approximately by a factor of 2. At the peak of the rainy season (September), the maximum
397 nocturnal R_{ch} values reached approximately $6.0 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ in FS and $9.0 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$
398 in Sh (Fig. 3, a and c). Thereafter, nocturnal R_{ch} declined well before the groundnut harvest along
399 with the rainfall spacing and the groundnut crop senescence (data not shown). During the dry
400 season nocturnal R_{ch} continued to decrease, with maximum values around $1.0 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$
401 in FS and $2.0 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ in Sh (Fig. 3, a and c).

402 The modeled nocturnal Rch values closely matched the measured nocturnal Rch values (mean
403 across four chambers per treatment), as indicated by the model performance metrics ($R^2 = 0.9$,
404 with bias and RMSE values of 0.3 and 0.5 $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$, respectively, for FS; $R^2 = 0.7$, with bias
405 and RMSE values of 0.4 and 0.6 $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$, respectively, for Sh) (Fig. 3, b and d). Similarly,
406 the daily mean modeled values also fitted well with the measured values, with FS showing (mean
407 + standard deviation) $0.9 \pm 0.9 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ (modeled) and $1.2 \pm 1.2 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$
408 (measured), while Sh recorded $1.4 \pm 0.9 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ (modeled) and $1.5 \pm 1.2 \mu\text{mol CO}_2 \text{ m}^{-2}$
409 s^{-1} (measured). Given the close match between the measured and modeled values, the fitted
410 model parameters were used subsequently to fill data gaps and estimate diurnal Rch values, as
411 presented in Fig. 4, a and c.



412 Fig. 3: Dynamics of instantaneous nocturnal CO₂ fluxes in chambers in Full sun (FS; a and b) and
 413 Shaded (Sh; c and d) environments (data filtered based on R² of the CO₂ variation over the time
 414 of chamber closure and Minimum Detectable Flux, Eq.2).

415 (a) and (c): measured nocturnal respiration in chambers (Rch: black dots; average of measurements in 4
 416 chambers per location) vs. modeled (coloured line). The vertical black line indicates the harvest date of
 417 groundnuts inside the chambers. The red arrows indicate the 'Birch' effect and the blue line represents the
 418 rainfall (mm 30mn⁻¹). Roman numerals (above the black arrows) refer to vegetation conditions prevailing
 419 inside the chambers, i.e. (I) bare soil, (II) weeds, (III) weeds + groundnuts, and (IV) groundnuts only.

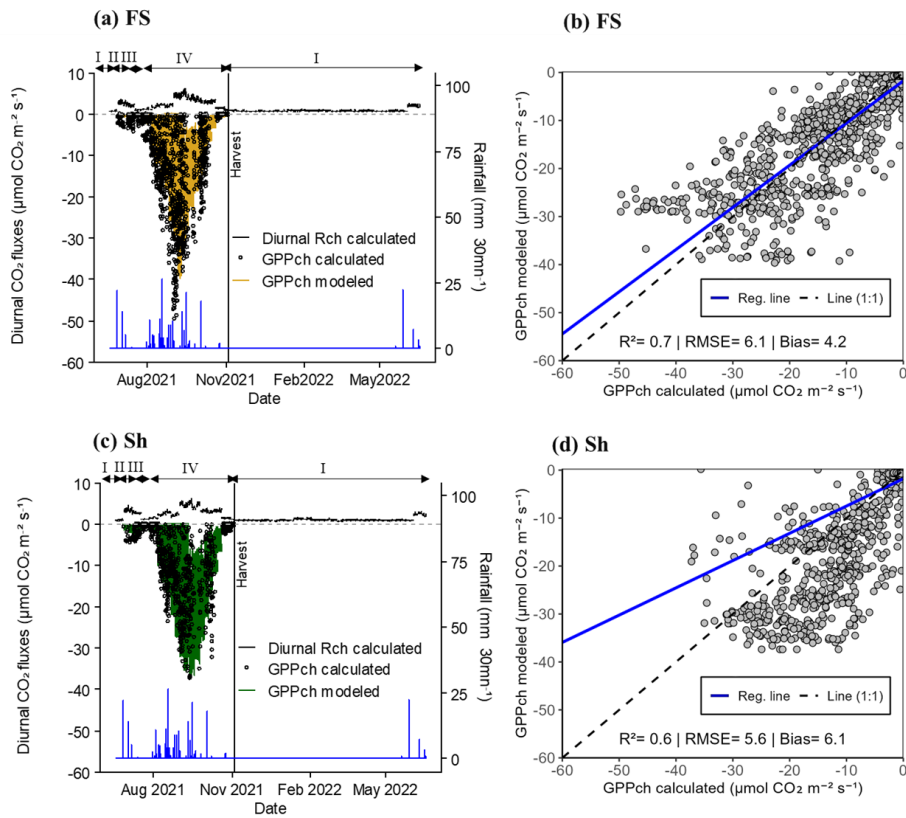
420 (b) and (d): scatter plot between measured and modeled nocturnal Rch. The solid blue line indicates the
 421 regression line and the dashed black one the (1:1) line RMSE and bias are expressed as fluxes (in μmol CO₂
 422 m⁻² s⁻¹). Each point represents the mean value from 4 chambers within the FS or Sh environments.

423 3.2.3. Dynamics of daytime fluxes in chambers

424 The measured GPPch.stand, as well as GPP modeled with Eq. 6, showed similar seasonal dynamics
425 inFS and Sh (Fig. 4, a and c). The fluxes peaked during the rainy season (Fig. 4a and c), coinciding
426 with periods of vigorous vegetative growth characterised by a high leaf area index (LAIch) of
427 groundnut plants within the chambers (S2, Fig. S32-4). The maximum calculated and standardised
428 GPPch values reached $-50 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ for FS and $-37 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ for Sh. As expected,
429 these fluxes were nil during the dry season when the soil was bare (Fig. 4, a and c).

430 The modeled GPPch values closely followed the same trends as the calculated values, although
431 model performance was slightly better for FS ($R^2 = 0.7$ with bias and RMSE values of 4.2 and 6.1
432 $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$, respectively) compared to Sh ($R^2 = 0.6$ with bias and RMSE values of 6.1 and
433 $5.6 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$, respectively) (Fig. 4, b and d).

434 The calculated diurnal respiration values (diurnal Rch calculated) for FS and Sh revealed a 'Birch
435 effect' similar to that observed during the night, though slightly more pronounced under Sh by a
436 factor of 1.2. Diurnal Rch values increased significantly during the rainy season, reaching a
437 maximum of $6.0 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ for both FS and Sh (Fig. 4, a and c). In the dry season, on bare
438 soil, these values declined, with maximum respiration reaching only $0.5 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ for both
439 situations (FS and Sh) (Fig. 4, a and c).



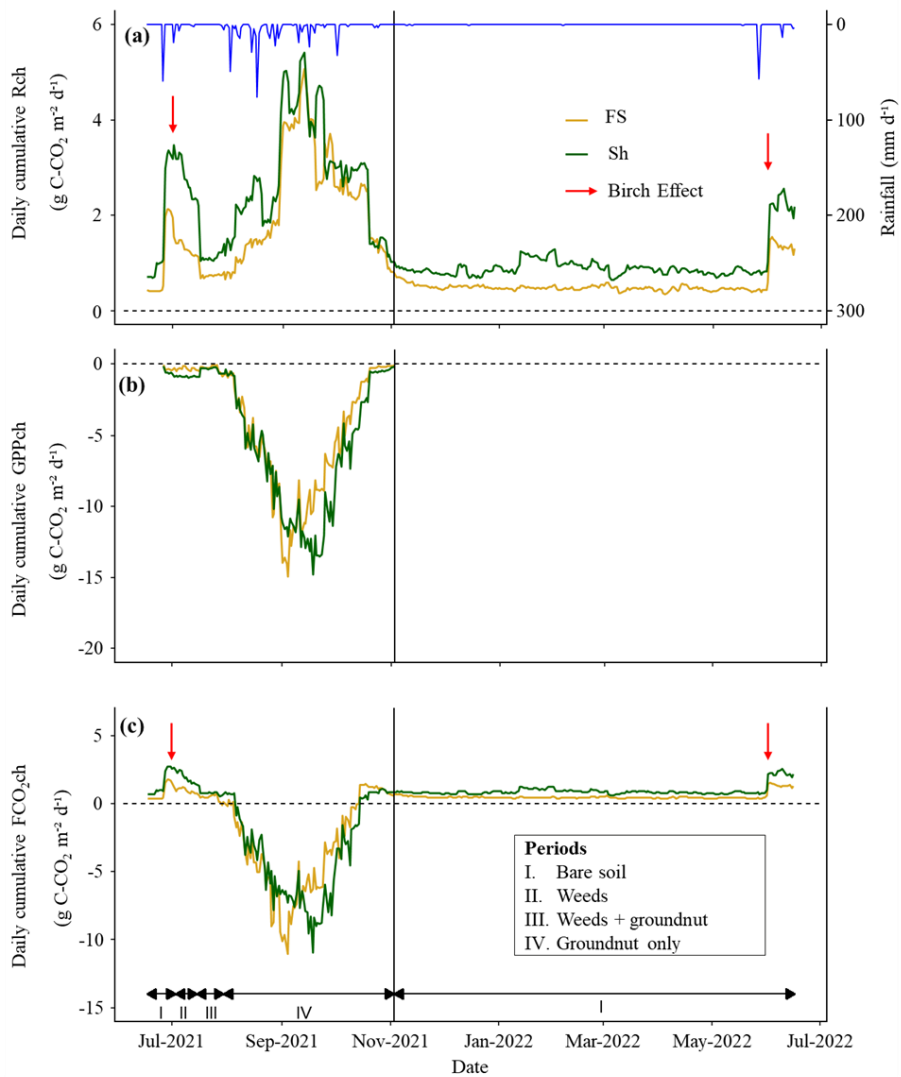
440 Fig. 4: Dynamics of instantaneous diurnal CO₂ fluxes in chambers in Full sun (FS; a and b) and
 441 Shaded (Sh; b and d) environments (filtered based on R² of the CO₂ variation over the time closure
 442 in FS and Sh and Minimum Detectable Flux, Eq.2).

443 (a) and (c): non-gap-filled diurnal Rch calculated (black line, positive values; average of measurements in
 444 4 chambers per location) and GPPch calculated from Eq.5 then standardised for LAI (black dots, negative
 445 values) and modeled (coloured line, negative values). The vertical black line indicates the harvest date of
 446 groundnuts inside the chambers and the blue line represents the rainfall (mm 30mn⁻¹). Roman numerals
 447 (above the black arrows) refer to conditions prevailing inside the chambers, i.e., (I) bare soil, (II) weeds,
 448 (III) weeds + groundnuts, and (IV) groundnuts.

449 (b) and (d): scatter plot between calculated and modeled GPPch. The solid blue line indicates the regression
 450 line and the dashed black one the (1:1) line. RMSE and bias are expressed as fluxes (in µmol CO₂ m⁻² s⁻¹).
 451 Each point represents the mean value from 4 chambers within the FS or Sh environments.

452 3.3. Dynamics of daily cumulative CO₂ fluxes in chambers

453 The seasonality of daily cumulative of GPPch.stand showed similar dynamics between FS and Sh,
454 with higher variability during the rainy season than during the dry season (Fig. 5). Daily total Rch
455 peaked during the rainy season at 5.1 g C-CO₂ m⁻² d⁻¹ for FS and 5.4 g C-CO₂ m⁻² d⁻¹ for Sh, while
456 the maximum GPPch.stand values were comparable at around -15.0 g C-CO₂ m⁻² d⁻¹ for both FS
457 and Sh (Table 1; ~~S2~~, Fig. ~~S72-4~~, a, b, c, and d). In the dry season, Rch decreased (Fig. 5), averaging
458 0.5 g C-CO₂ m⁻² d⁻¹ for FS and 1.0 g C-CO₂ m⁻² d⁻¹ for Sh. GPPch declined well before harvest
459 (senescence) and remained nil during the dry season (Fig. 5). During the rainy season FCO₂ch
460 peaked in absolute value at around 11.0 g C-CO₂ m⁻² d⁻¹ for FS and Sh (Fig. 5) (Table 1; ~~S2~~, Fig.
461 ~~S72-4~~, e and f), while FCO₂ch values were the same as Rch during the dry season. In absolute terms,
462 the mean Rch and GPPch were significantly higher under Sh as compared to FS, by factors of 1.3
463 and 1.2, respectively. Conversely, the mean FCO₂ch was significantly higher in absolute value
464 under FS (0.4 g C-CO₂ m⁻² d⁻¹) than under Sh (0.2 g C-CO₂ m⁻² d⁻¹) (Table 1).
465 The annual cumulative Rch values were 392.8 g C-CO₂ m⁻² for FS and 574.5 g C-CO₂ m⁻² for Sh.
466 The GPPch fluxes reached -539.5 g C-CO₂ m⁻² for FS and -632.6 g C-CO₂ m⁻² for Sh. The net Annual
467 ~~net~~ cumulative C exchange (FCO₂ch) ~~were~~was -146.7 g C-CO₂ m⁻² in FS and -58.1 g C-CO₂ m⁻² in
468 Sh.



469 Fig. 5: Seasonal dynamics of daily gap-filled cumulative fluxes (in $\text{gC-CO}_2 \text{ m}^{-2} \text{ d}^{-1}$) in chambers.

470 (a) soil+crop respiration (Rch), (b) photosynthesis (GPPch, standardised for LAI) and (c) net CO_2 exchange
 471 ($\text{FCO}_{2\text{ch}}$). The yellow and green solid lines compare the FS and Sh environments, respectively. The vertical
 472 black line indicates the harvest date of groundnuts inside the chambers. The blue line depicts the daily
 473 cumulative rainfall (mm d^{-1}) throughout the rainy season, and the red arrow indicates the 'Birch'
 474 effect. Roman numerals (above the black arrows) in (a) and (c) refer to the prevailing conditions inside the
 475 chambers: (I) bare soil, (II) weeds, (III) weeds + groundnuts, (IV) groundnuts.

476 Table 1: Comparison of daily cumulative and gap-filled chamber CO₂ fluxes (Rch, GPPch
 477 standardised for LAI, and FCO₂ch in g C-CO₂ m⁻²) in the FS and Sh condition.
 478

	Annual sum	Daily Mean ±SD	Min	Max	Mann-Whitney test
(g C-CO ₂ m ⁻²)	.yr ⁻¹	.d ⁻¹	.d ⁻¹	.d ⁻¹	
Rch					
FS	392.8	1.1 ± 0.9	0.4	5.1	*
Sh	574.5	1.6 ± 1.1	0.6	5.4	
GPPch					
FS	-539.5	-4.1 ± 4.3	< -0.1	-14.9	*
Sh	-632.6	-4.8 ± 4.6	< -0.1	-14.8	
FCO₂ch					
FS	-146.7	-0.4 ± 2.4	-11.0	1.8	*
Sh	-58.1	-0.2 ± 2.7	-10.9	2.8	

479 Annual sum corresponds to the annual cumulative fluxes (g C-CO₂ m⁻² yr⁻¹). Mean, SD, Min, and Max
 480 represent respectively the mean, standard deviation, minimum, and maximum values at the daily scale (g
 481 C-CO₂ m⁻² d⁻¹). Asterisks (*) indicate the p-values from the Mann-Whitney test, used to assess differences in
 482 mean between FS and Sh (p < 0.05). Positive values indicate CO₂ emissions, while negative values represent
 483 CO₂ uptake.

484 3.4. Drivers of daily respiration and photosynthesis in chambers

485 The chamber-based daily cumulative respiration (Rch) and GPPch showed significant and positive
486 correlations with the leaf area index (LAIch), both at a distance from the trees (FS) and under the
487 trees (Sh) (Table 2). The influence of LAIch on GPPch was stronger ($r = 0.86$ for FS and Sh) than
488 its influence on Rch ($r = 0.60$ for FS; $r = 0.69$ for Sh). Soil VWC was also positively correlated with
489 Rch and GPPch, both in FS and Sh. However, the influence of soil VWC on Rch was stronger under
490 Sh compared to FS, while its influence on GPPch was similar in both situations (FS and Sh). Soil
491 temperature showed weak negative correlations with Rch (in FS and Sh) and with GPPch (only in
492 Sh). Finally, no significant correlations were found between T_{air} , and any of the CO_2 fluxes (Table
493 2).

494 Table 2: Spearman correlation matrix based on daily cumulative and gap-filled CO₂ fluxes from full
 495 year chamber measurements (g C-CO₂ m⁻² d⁻¹) with microclimatic parameters.

Parameters	Condition	r-coef. Rch	p (Rch)	r-coef. GPPch	p (GPPch)
T _{soil}	FS	-0.25 ***	<u>7.47 x 10⁻⁴</u>	ns	<u>1.18 x 10⁻³</u>
	Sh	-0.28 ***	<u>9.69 x 10⁻¹⁴</u>	-0.38 ***	<u>2.88 x 10⁻⁷</u>
T _{air}	FS	ns	<u>0.22</u>	ns	<u>0.35</u>
	Sh	ns	<u>0.98</u>	ns	<u>0.15</u>
VWC	FS	0.51 ***	<u>3.00 x 10⁻³⁴</u>	0.75 ***	<u>6.73 x 10⁻³</u>
	Sh	0.78 ***	<u>1.29 x 10⁻⁶⁶</u>	0.75 ***	<u>0.02</u>
LAIch	FS	0.60 ***	<u>1.11 x 10⁻⁶¹</u>	0.86 ***	<u>2.23 x 10⁻⁸</u>
	Sh	0.69 ***	<u>6.08 x 10⁻⁶⁹</u>	0.86 ***	<u>2.11 x 10⁻¹²</u>

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496 Spearman correlation coefficients (r-coef) between daily cumulative and gap-filled CO₂ flux components
 497 (Rch and GPPch, with GPPch in absolute terms) and daily mean microclimatic parameters in full sun (FS)
 498 and shaded chambers (Sh). T_{soil} (°C) is the daily mean soil temperature at 6 cm depth, T_{air} (°C) the daily
 499 mean air temperature at 15 cm height, VWC (%) the daily mean volumetric water content (VWC, %), and
 500 LAIch (m⁻² leaf m⁻² soil) the chamber leaf area index value for a given day. Letter p represents the p-value
 501 and significance levels are indicated by (***) p < 0.001; ** p < 0.01; * p < 0.05) for p < 0.001; ns denotes a non-
 502 significant correlation (p > 0.05).

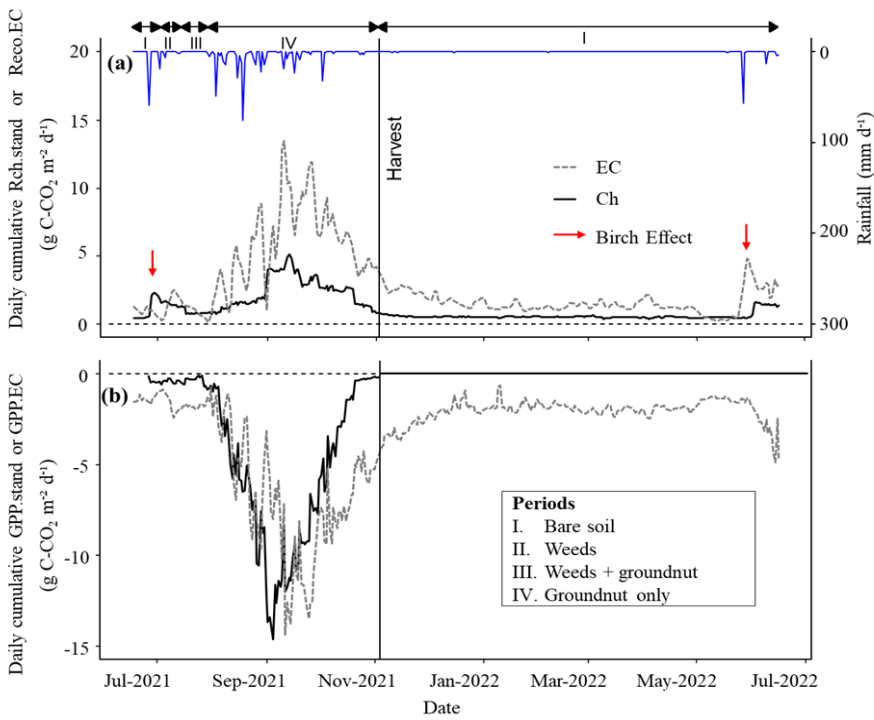
503 3.5. Comparison of respiration and GPP measurements between chambers (Ch) and Eddy
504 Covariance (EC) methods

505 The chamber-based daily total CO₂ fluxes, gap-filled and weighted according tree cover were
506 compared with the fluxes obtained using the EC method (Fig. 6).

507 During the rainy season, both total respiration and GPP showed comparable dynamics between
508 the two methods, with synchronised peaks and higher variability compared to the dry season (Fig.
509 6). The maximum value of Reco.EC, peaked at 13.5 g C-CO₂ m⁻² d⁻¹ (Table 3). The initial value of
510 Rch.stand was comparable to Reco.EC but peaked only at 5.1 g C-CO₂ m⁻² d⁻¹ (Table 3), meaning
511 a third of the peak of Reco.EC. The maximum GPP, was -14.3 g C-CO₂ m⁻² d⁻¹ and -14.6 g C-CO₂
512 m⁻² d⁻¹ for GPP.EC and GPPch.stand, respectively (Table 3). This indicates that the LAI-based
513 standardisation and upscaling approach were realistic, at least up to the peak of groundnut
514 growth.

515 On average, Reco.EC was significantly higher than Rch.stand, by a factor of 2.3. GPP.EC was also
516 significantly higher than GPPch.stand, but only by a factor of 1.2 (Table 3).

517 During the dry season, Reco.EC and Rch.stand gradually decreased. The values for Reco.EC
518 remained higher than for Rch.stand, which was fairly consistent with the contribution of the Ra
519 tree above-ground compartment, even if this difference seemed to disappear at the end of the dry
520 season (Fig. 6). The measured 'Birch effect' was highest for Rch.stand in 2021, but was the
521 opposite in 2022 due to a system failure at the beginning of the rainy season. The maximum value
522 of GPP.EC reached -2.4 g C-CO₂ m⁻² d⁻¹ when the trees were at their maximum of foliage, after
523 harvest and while weeds were still present in the field. However, after the harvest, chamber
524 photosynthesis (GPPch.stand) was nil (Table 3).



525 Fig 6: Comparing the seasonal dynamics of CO₂ fluxes between Eddy Covariance (EC)
 526 measurements and upscaled chamber measurements (ch.stand).

527 (a) represent the seasonal dynamics of soil + crop respiration (Rch.stand) and ecosystem respiration
 528 (Reco.EC) and (b) photosynthesis (GPP.stand and GPP.EC). The black and dashed grey lines show Ch and
 529 EC seasonal dynamics, respectively. The vertical black line indicates the harvest date of groundnuts inside
 530 the chambers. The blue line depicts the daily cumulative rainfall (mm d⁻¹), and the red arrow indicates the
 531 'Birch' effect. Roman numerals (above the black arrows) refer to conditions prevailing inside the
 532 chambers: (I) bare soil, (II) weeds, (III) weeds + groundnuts, (IV) groundnuts.

533 Table 3: Comparison of gap-filled CO₂ fluxes between Eddy Covariance (EC) and upscaled chamber (Ch.stand) measurements, by season (rainy or dry).

	Rainy season				Dry season			
	Daily Mean ± SD	Min	Max	Mann-Whitney test	Daily Mean ± SD	Min	Max	Mann-Whitney test
(g C-CO ₂ m ⁻²)	.d ⁻¹	.d ⁻¹	.d ⁻¹		.d ⁻¹	.d ⁻¹	.d ⁻¹	
Reco.EC or Rch.stand								
EC	4.6 ± 3.2	0.2	13.5	*	1.2 ± 0.4	0.3	2.1	*
Ch.stand	2.0 ± 1.1	0.5	5.1		0.5 ± 0.04	0.4	0.6	
GPP.EC or GPPch.stand								
EC	-5.1 ± 3.6	-0.7	-14.3	*	-1.7 ± 0.3	-0.6	-2.4	
Ch.stand	-4.2 ± 4.3	<-0.1	-14.6		0	0	0	

534 Mean, SD, Min, and Max represent the daily mean fluxes, standard deviation, minimum, and maximum values, respectively (g C- CO₂ m⁻² d⁻¹). The Asterisks (*) indicate
535 the p-values from the Mann-Whitney test, used to assess differences in mean between EC and Ch. Positive values indicate CO₂ emissions, while negative values
536 represent CO₂ uptake.

537 *3.6. The contribution of F. albida to Reco and GPP*

538 During the dry season, the cumulative contribution of *F. albida* to ecosystem respiration (Ra tree)
539 was 139.6 g C-CO₂ m⁻². This represent ~~14%~~~~12%~~ of the total annual cumulative Reco, which was
540 estimated at ~~1000.0~~~~1180.0~~ g C-CO₂ m⁻² (Table S4). The contribution of trees (GPP tree) to total
541 annual GPP in absolute term was -270.2 g C-CO₂ m⁻², equivalent to ~~~23%~~~~50%~~ of the total annual
542 cumulative GPP of the ecosystem measured by EC (~~1180.0~~~~550~~ g C-CO₂ m⁻²) (Table S4).
543 The ratio between these two components (Ra tree / GPP tree) in absolute terms was 0.52,
544 reflecting a carbon use efficiency (CUE) of 0.48 (~~S3~~, Table S~~43.4~~).

545 *3.7. Annual vertical CO₂ balances* ~~Carbon budgets~~ *at the field-scale*

546 The upscaled chamber-based annual cumulative total respiration flux (Rch.stand) was estimated
547 to be 4.1 ± 0.~~1804~~ Mg C-CO₂ ha⁻¹ (Table 4). In comparison, the annual ~~budget of~~ Reco.EC was 10.0
548 ± 0.~~4903~~ Mg C-CO₂ ha⁻¹ (Table 4), more than two times larger than Rch.stand.
549 The upscaled GPPch.stand reached an annual cumulative value of -5.5 ± 0.~~8303~~ Mg C-CO₂ ha⁻¹,
550 whereas the annual cumulative GPP.EC was -11.8 ± 0.~~5303~~ Mg C-CO₂ ha⁻¹ (Table 4).
551 The net annual vertical net-C budget~~balance~~, based on both methods, was estimated at -1.4 ±
552 0.~~4602~~ Mg C-CO₂ ha⁻¹ for chambers (FCO₂ch.stand) and -1.8 ± 0.~~1704~~ Mg C-CO₂ ha⁻¹ for Eddy
553 Covariance (NEE.EC) (Table 4).

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554 Table 4: Annual budget of CO₂ fluxes based on Eddy Covariance (EC) and upscaled chamber
 555 methods (Ch.stand).

	Annual sum (Mg C-CO ₂ ha ⁻¹)	Std error (Mg C-CO ₂ ha ⁻¹)
Reco.EC or Rch.stand		
EC	10.0	0.4903
Ch.stand	4.1	0.1801
GPP.EC or GPPch.stand		
EC	-11.8	0.5303
Ch.stand	-5.5	0.8303
NEE.EC or FCO₂ch.stand		
EC	-1.8	0.1701
Ch.stand	-1.4	0.4602

556 Annual sum corresponds to the annual cumulative fluxes for full year measurements (Mg C-CO₂ ha⁻¹). EC
 557 refers to fluxes measured by the Eddy Covariance method, and Ch refers to the fluxes measured by
 558 chambers, which are then upscaled to the whole field. Rch.stand represents the chamber respiration, while
 559 Reco.EC denotes the ecosystem respiration according to the EC method. GPP.EC and GPPch.stand are the
 560 gross primary production or photosynthesis flux, measured by EC and Ch methods, respectively. NEE.EC
 561 and FCO₂ch.stand represent the net ecosystem exchange for EC and Ch, respectively. The associated
 562 standard error is denoted as Std error (Mg C-CO₂ ha⁻¹). Positive values indicate CO₂ emissions, while
 563 negative values represent CO₂ uptake.

564 4. Discussion

565 4.1. Soil respiration modeling and limitations regarding the temperature

566 In this study the Lloyd and Taylor (1994) model, based on a modified Arrhenius-type formulation,
567 was used to model nocturnal soil respiration fluxes for estimating daytime respiration. Unlike the
568 classical Arrhenius equation, this model includes the $(T_{\text{soil}} - T_0)$ term in the denominator of the
569 exponential expression (Eq. 4), which inherently limits the effects of high temperatures by
570 progressively reducing the temperature sensitivity of soil respiration as temperatures rise above
571 a given threshold. This structural feature produces a flattening of the respiration-temperature
572 relationship at elevated temperatures, thereby preventing the overestimations (Lloyd and Taylor,
573 1994).

574 The Lloyd and Taylor model has successfully been widely applied, primarily in boreal and
575 temperate ecosystems (Lasslop et al., 2010; Reichstein et al., 2003), and relies on the assumption
576 of comparable thermal conditions between daytime and nighttime periods (Juszczak et al., 2012).
577 In our study, instantaneous soil temperatures ranged from 20.7 to 45.8 °C during the day and from
578 22.1 to 45.0 °C at night, indicating largely overlapping thermal ranges between the two periods.
579 Model parameters were recalibrated using five-day fixed windows, which provided sufficient
580 temporal resolution while capturing seasonal dynamics of soil respiration.

581 This study represents one of the first applications of the Lloyd and Taylor model in a Sahelian
582 semi-arid context. While Arrhenius based models are known to potentially overestimate fluxes
583 under extreme temperatures due to physiological limitations, over the range of temperatures
584 observed in this study, the modeled soil respiration was not overestimated (Fig. S6, a and b). Thus,
585 the model used in this study appears to provide a realistic representation of soil respiration under
586 local conditions. However, this conclusion is site-specific and should not be interpreted as a
587 general validation of temperature-based models across all semi-arid environments. Such models
588 should be systematically validated with respect to temperature to ensure their reliability.

589 4.1.4.2. Seasonality and drivers of chamber-based CO₂ fluxes

590 In our agroforestry context, seasonal variability in CO₂ fluxes closely followed rainfall dynamics,
591 peaking during the wet season and declining sharply in the dry season, consistent with soil
592 moisture depletion and crop senescence. This pattern is typical of semi-arid ecosystems (Ago et
593 al., 2016a; Brümmer et al., 2008; Guillen-Cruz et al., 2023; Macharia et al., 2020; Mosongo et al.,
594 2022; Wieckowski et al., 2024).

595 Respiration and photosynthesis were primarily driven by soil moisture and LAI, reflecting the
596 system's sensitivity to water availability and crop dynamics. Soil moisture enhanced both
597 processes by stimulating microbial activity and supporting plant growth (Borken et al., 2002;

598 Conant et al., 2004; Merbold et al., 2009; Yu et al., 2020; Zhao et al., 2016). The stronger correlation
599 between soil moisture and respiration under *F. albida* canopy (Sh: $r = 0.78$) compared to
600 full sun (FS: $r = 0.51$) suggests greater microbial sensitivity to moisture beneath trees. This likely
601 reflects enhanced substrate availability, resulting in stronger post-rainfall respiration pulses
602 (Meisner et al., 2015) and supporting the 'fertile island' effect, where trees improve local soil
603 conditions (Eldridge et al., 2024). Photosynthetic capacity also responded to soil moisture, as
604 shown by positive correlations with LAI and key physiological traits such as light use efficiency
605 (α) and maximum CO₂ uptake rate (β) (Gonsamo et al., 2019; Qiu et al., 2023; Zhang et al., 2024).
606 In contrast, the influence of soil temperature (T_{soil}) on respiration was weakly negative in both FS
607 and Sh, indicating a thermal threshold beyond which respiration is suppressed—estimated at 32
608 ± 1.5 °C in FS and 29.5 ± 1.9 °C in Sh (S2, Fig. S2.6, a and b), similar to findings in Eastern Ghana
609 (Owusu et al., 2024). This inhibition likely results from decreased enzymatic and microbial
610 activity under combined heat and water stress (Liu et al., 2018; Richardson et al., 2012). In semi-
611 arid regions, soil respiration often becomes decoupled from temperature due to seasonal
612 moisture constraints (Jia et al., 2020; Tucker & Reed, 2016; Warren, 2014), with microbial activity
613 limited during dry periods despite favourable temperatures. This decoupling helps explain the
614 weak or absent correlation between T_{soil} and soil moisture (S2, Fig. S2.5, b), particularly under
615 Sh ($r = -0.28$). Management practices such as organic inputs can also modulate these dynamics,
616 adding further variability to soil respiration responses (Meena et al., 2020; Oyonarte et al., 2012;
617 Rong et al., 2015; Xue & Tang, 2018).

618 4.2.4.3. Magnitude of chamber-based total CO₂ respiration fluxes

619 Mean total soil respiration values were consistent with those reported in other low-input
620 agricultural systems across sub-Saharan Africa (Mapanda et al., 2010; Pelster et al., 2017;
621 Rosenstock et al., 2016). In full sun (FS), the mean respiration (1.0 ± 0.9 g C-CO₂ m⁻² d⁻¹) closely
622 matched values measured by Wachiye et al. (2020) in a semi-arid Kenyan field at 1158 m altitude
623 (1.1 ± 0.1 g C-CO₂ m⁻² d⁻¹). This similarity likely reflects comparable environmental conditions,
624 including moderate rainfall (~ 550 mm yr⁻¹) and low soil organic carbon and nitrogen contents
625 ($<1\%$) in the 0–20 cm layer of sandy soil. In contrast, respiration under *F. albida* canopy (Sh: 1.6
626 ± 1.1 g C-CO₂ m⁻² d⁻¹) was higher, likely due to additional autotrophic respiration from tree roots
627 and greater organic inputs beneath the canopy. Nonetheless, this flux remains close to values
628 observed in low-input sorghum fields on sandy loam soils in eastern Ghana (1.7 ± 1.1 g C-CO₂ m⁻²
629 d⁻¹), despite higher rainfall (950–1000 mm yr⁻¹) in that region (Owusu et al., 2024).
630 Cumulative annual respiration fluxes fell within the range reported for Sahelian croplands (250–
631 450 g C-CO₂ m⁻²) (Brümmer et al., 2009) and other sub-Saharan African agricultural systems (Kim
632 et al., 2016). The cumulative flux under tree cover is similar to that measured in cassava fields in

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633 eastern Tanzania ($440 \text{ g C-CO}_2 \text{ m}^{-2} \text{ yr}^{-1}$), despite the latter receiving higher rainfall ($\sim 1115 \text{ mm}$
634 yr^{-1}) (Rosenstock et al., 2016). This convergence may stem from comparable soil fertility
635 constraints, with low soil organic carbon (1–1.7%) and nitrogen contents ($<0.5\%$). In contrast,
636 the slightly lower cumulative flux in FS may reflect less favourable microclimatic conditions—
637 such as elevated soil temperatures and increased aridity away from tree cover—limiting
638 microbial activity (see Section 4.1).

639 Across sub-Saharan Africa, soil respiration fluxes based on static chamber measurements show
640 high spatial variability, largely shaped by climate and land use. For example, Owusu et al. (2024)
641 found higher respiration in woodlands ($3.8 \pm 0.8 \text{ g C-CO}_2 \text{ m}^{-2} \text{ d}^{-1}$) and grazed areas (2.7 ± 1.7)
642 than in croplands (1.7 ± 1.1) in humid eastern Ghana. This gradient was linked to differences in
643 soil moisture and organic matter. Similarly, Rosenstock et al. (2016) reported much higher fluxes
644 in highland pastures in Kenya ($3.8\text{--}4.4 \text{ g C-CO}_2 \text{ m}^{-2} \text{ d}^{-1}$) compared to cultivated fields in eastern
645 Tanzania (1.2 ± 0.2), highlighting the role of vegetation cover and soil fertility.

646 [4.3.4.4](#) *Effect of trees on chamber-based soil respiration and photosynthesis*

647 A notable increase in respiration and photosynthesis fluxes was observed under *F. albida* trees
648 (Sh) compared at a distance from trees (FS). This increase may indicate the potential role of *F.*
649 *albida* in modulating CO_2 exchange dynamics (Rch and GPPch) within this agro-silvo-pastoral
650 system. These results are consistent with preliminary findings from similar environments
651 (Duthoit et al., 2020).

652 Numerous studies have investigated the effect of tree species on greenhouse gas fluxes,
653 particularly CO_2 , revealing significant variations across different ecological contexts (Bréchet et
654 al., 2021, 2025; Klaus et al., 2024; Mazza et al., 2021; Ramesh et al., 2013; Rheault et al., 2024).
655 However, the underlying mechanisms by which trees influence these dynamics are not yet fully
656 understood.

657 In general, agroforestry systems have been well-documented for their ability to provide a range
658 of ecosystem services (e.g., Assefa et al., 2024; Bado et al., 2021; Kuyah et al., 2019; Rolo et al.,
659 2023). Specifically, *Faidherbia*-based agroforestry systems may play a crucial role in regulating
660 CO_2 exchanges between the soil and atmosphere. *F. albida*-based agroforestry systems are
661 recognized for enhancing both soil organic and mineral fertility (Bayala et al., 2020; Dilla et al.,
662 2019; Sileshi, 2016; Sileshi et al., 2020; Stephen et al., 2020), mainly through litter accumulation
663 and direct inputs from livestock excreta under their canopies. Additionally, the extensive roots
664 system of *F. albida* trees helps concentrate mineral nutrients, contributing to the formation of a
665 'fertile island' effect under the trees (Siegwart et al., 2022; Eldridge et al., 2024). Moreover, *F.*
666 *albida* improve water infiltration (Diongue et al., 2023; Faye et al., 2020; Sarr et al., 2023), enhance
667 soil moisture retention (Clermont-Dauphin et al., 2023) and contribute to reduced soil

668 temperatures (de Carvalho et al., 2021; Lopes et al., 2024; Sida et al., 2018). These changes foster
669 a more favourable environment for soil microbial activity and crop development (Diack et al.,
670 2024; Diene et al., 2024; Leroux et al., 2020; Roupsard et al., 2020) under the trees compared to
671 open areas. Consequently, this likely explains the stronger effect of soil moisture and the leaf area
672 index of groundnuts on Rch under the trees, resulting in higher total respiration (Table 2). For
673 photosynthesis, the effect of these parameters was similar in both FS and Sh (Table 2). However,
674 the significantly higher intensity of GPPch under Sh can be explained by greater light use efficiency
675 (α) and a higher maximum CO₂ uptake rate at light saturation (β) in this shaded environment. In
676 agroforestry systems, light use efficiency can at least partially mitigate the reduction in
677 photosynthetically active radiation under tree canopies (Charbonnier et al., 2017).
678 Similar results have been observed in different climatic conditions and ecosystems. Gomes et al.
679 (2016) investigated soil respiration using mobile chambers (LI-8100-102 model) under trees in
680 coffee-based agroforestry (AF) systems and in the open areas (FS) in Minas Gerais, Brazil. These
681 studies were conducted with agroecological management practices, such as weeding,
682 intercropping maize between coffee rows, and mulching. The AF-agroforestry systems exhibited
683 lower air and soil temperatures (at 5 and 10 cm depth) and higher air and soil humidity compared
684 to the open areas FS (Gomes et al., 2016). These authors observed greater spatial variability in soil
685 respiration in agroforestry system AF (34.1%) compared to the open areas FS (24.2%). This
686 variability was mainly linked with fluctuations in labile carbon and total nitrogen, reflecting more
687 favourable soil microclimate for microbial activity in agroforestry system AF. In contrast, soil
688 temperature (10 cm depth) accounted for most of the variability observed in the open areas FS,
689 where the absence of tree canopy resulted in high soil temperatures and low soil moisture (Gomes
690 et al., 2016). Likewise, Haren et al. (2010) reported 38% higher soil respiration near large trees
691 (DBH > 35 cm) in clay-rich Amazonian forests compared to open sites. Interestingly, the
692 magnitude of CO₂ fluxes was independent of tree species, indicating that canopy effects may
693 outweigh species-specific traits in some contexts. In our study, *F. albida*'s influence on CO₂ fluxes
694 aligns with this general pattern observed in tropical agroforestry. However, the mechanisms
695 linking individual tree species to microbial and physicochemical drivers of CO₂ dynamics remain
696 insufficiently understood and warrant further investigation (Jevon et al., 2023).

697 4.4.4.5. Birch Effect

698 A rapid increase in soil respiration was observed following the first rainfall events, particularly
699 under *F. albida*. This phenomenon can be attributed to the lower bulk density of the soil under the
700 trees (Clermont-Dauphin et al., 2023; Siegwart et al., 2023), which potentially lead to CO₂
701 accumulation during the dry season due to higher soil organic matter (SOM) (Siegwart et al.,
702 2023). Additionally, the sensitivity of microbial communities to subtle variations in soil moisture,

703 compounded by the tree effect, may further explain this phenomenon, as outlined in Sections 4.1
704 and 4.3. This phenomenon, known as the 'Birch effect' (Birch, 1958), has been reported across
705 various semi-arid ecosystems in sub-Saharan Africa (Ago et al., 2016b; Fan et al., 2015;
706 Wieckowski et al., 2024), as well as other semi-arid ecosystems globally (Roby et al., 2022; Yan et
707 al., 2014; Yu et al., 2020). In these contexts, the 'Birch effect' may result from the displacement of
708 soil gas phases by the piston effect generated during water infiltration (Singh et al., 2023).
709 Furthermore, microbial communities in semi-arid environments adopt osmoregulatory
710 mechanisms to withstand water deficit (Warren, 2014), which is particularly pronounced during
711 the dry season. This phenomenon reduces soil microbial metabolism (Schimel et al., 2007). Upon
712 rapid soil rewetting, especially after prolonged dry periods, soil microbial metabolism process is
713 swiftly reactivated, leading to a transient pulse in respiration and a CO₂ release (Barnard et al.,
714 2020; Kim et al., 2012; Manzoni et al., 2020; Vargas et al., 2018). Isotopic signatures of soil
715 respiration provide evidence supporting the hypothesis that these pulses result from the rapid
716 mineralisation of necromass or osmolytes excreted by microorganisms under drought stress
717 (Schimel et al., 2007; Unger et al., 2010). Additional factors may amplify the 'Birch effect'. For
718 instance, drying-rewetting cycles can induce physical disruption of soil aggregates, enhance
719 oxygen penetration and thereby expose previously protected organic matter to microbial
720 decomposition (Rabbi et al., 2024). This increases substrate availability and subsequently boosts
721 soil respiration fluxes.
722 The magnitude of the 'Birch effect' is modulated by the severity and duration of drought. Thus, at
723 our study site, given the 8- to 9-month-long dry season, the 'Birch effect' is particularly intense.
724 Indeed, extended drought periods promote greater accumulation of microbial necromass and
725 intensify hypo-osmotic stress responses upon rewetting (Singh et al., 2023).

726 [4.5.4.6](#) *Comparing EC and chamber-based methods*

727 Results revealed high seasonal variability, with higher values during the rainy season compared
728 to the dry season. This seasonal pattern aligns with findings from studies in the Sahel using the
729 EC method for CO₂ flux measurements (Brümmer et al., 2008; Tagesson et al., 2015; Agbohessou
730 et al., 2023, Wieckowski et al., 2024). Comparable patterns have ~~also~~ been ~~also~~ documented at the
731 ecosystem scale in other semi-arid environments (Ago et al., 2014; Archibald et al., 2009; Ardö et
732 al., 2008; Jia et al., 2020; Quansah et al., 2015; Williams et al., 2009; Zhang, Bi, et al., 2024).
733 Several comparative studies between chamber and EC methods have reported both congruent
734 and divergent CO₂ flux estimates (Bastviken et al., 2022; Poyda et al., 2017; Riederer et al., 2014;
735 J. Tang et al., 2008; Wang et al., 2010). In the present study, ecosystem respiration fluxes during
736 the rainy season exhibited notable discrepancies measurements between EC (Reco.EC) and
737 upscaled chamber-based (Rch.stand). This is attributable to differences in the flux components

738 captured by each method. Specifically, Reco.EC included respiration from below- and above-
739 ground tree parts, crops (groundnuts and cowpeas), weeds, and soil, whereas Rch.stand
740 accounted only respiration from below-ground tree, groundnut crop, and soil. Therefore, as
741 expected, Reco.EC (4.6 ± 3.2 g C-CO₂ m⁻² d⁻¹) were significantly higher than Rch.stand (2.0 ± 1.1 g
742 C-CO₂ m⁻² d⁻¹).

743 For chamber-based GPP measurements, values were standardised (GPP-stand) by the field's leaf
744 area index (LAI.field). This allowed it to improve comparability with GPP.EC when trees were
745 leafless in the rainy season. In both cases, GPP accounted only for crops (groundnut and cowpea)
746 and weeds, as trees were non-photosynthetic in the rainy season. Despite this standardisation,
747 GPP.EC values (-5.1 ± 3.6 g C-CO₂ m⁻² d⁻¹) were significantly higher than GPPch.stand values (-4.2
748 ± 4.3 g C-CO₂ m⁻² d⁻¹). However, ~~the no divergence~~ **was observed in August, and the intensity of**
749 **the peak of GPP in September was similar in both methods**~~did not occur on the peak of GPP (which~~
750 ~~was very similar in both methods)~~, but from the onset of groundnut senescence, when weeds
751 became the dominant photosynthetic contributors. Thus, during the groundnut growth season,
752 with leafless *F. albida* trees and almost no weeds, GPP measurements from EC and chambers
753 generate closely comparable results. Therefore, this provides an initial form of cross-validation
754 between the two methods. It is important to note that the EC method integrates CO₂ fluxes over a
755 larger spatial scale, encompassing all ecosystem components (Baldocchi, 2003), while the
756 chamber method captures fluxes on a smaller scale (i.e., at the 0.25 m² scale). This scale disparity
757 can introduce uncertainties when upscaling chamber-based fluxes to the field, as vegetation
758 composition within chambers does not represent the EC footprint's average vegetation. This
759 makes upscaling chamber-based measurements challenging. Nevertheless, the standardisation
760 we applied on chamber photosynthesis by LAI has been relatively successful.

761 During the dry season, Reco.EC included respiration from below- and above-ground tree parts
762 (with leaves) and bare soil, while Rch.stand measured only below-ground tree and bare soil
763 respiration. Consequently, the difference between Reco.EC and Rch.stand was solely attributable
764 to above-ground tree respiration (R_a tree above-ground). In terms of GPP, chamber
765 measurements were nil, whereas GPP.EC reflected only GPP trees.

766 The transition period, characterised by groundnut senescence, tree leaf regrowth, and weed
767 proliferation, introduced further complexity, amplifying method-specific discrepancies. Rch.stand
768 measurements facilitated the estimation of tree contribution to Reco.EC (R_a tree) and the
769 verification of the consistency for EC results in terms of carbon use efficiency (CUE), estimated
770 here at 0.48. This value indicates that nearly 50% of the carbon captured by trees is allocated to
771 biomass. The CUE estimate here is well comparable to the global average across diverse
772 ecosystems, climates, and management practices (0.49 ± 0.14) (Tang et al., 2019). Similar CUE
773 values have been reported for semi-arid grasslands (0.46 ± 0.10), but our value is notably lower

774 than those documented for wetlands (0.61 ± 0.13) (Tang et al., 2019). Overall, these findings
775 reinforce the plausibility of our assumptions regarding the compartment-~~s~~ contributions to
776 Reco.EC and Rch.stand, thereby providing a second cross-validation of the EC-Ch comparison.
777 However, despite a frequently assumed CUE of 0.5 in models, global estimates span a broad range
778 (0.20 to 0.82), depending on ecosystem type and management practices (DeLucia et al., 2007;
779 Tang et al., 2019). This underscores the importance of refining carbon flux models to better
780 represent the biophysical processes governing CO₂ exchange in semi-arid agroforestry systems.
781 The combined use of EC and chamber methodologies offers a comprehensive perspective on
782 ecosystem-scale CO₂ flux dynamics, advancing the understanding of carbon cycling in these
783 environments.

784 *4.6.4.7. Net annual vertical carbon balance ~~carbon exchange budget~~*

785 The net annual vertical carbon balance ~~annual net carbon (C) exchange budget~~ was quantified at
786 -1.4 ± 0.4602 Mg C-CO₂ ha⁻¹ with the chamber method and -1.8 ± 0.1701 Mg C-CO₂ ha⁻¹ by the
787 Eddy Covariance (EC), indicating that the studied agro-silvo-pastoral system functions as a net
788 carbon sink. These findings corroborate the system-~~s~~ potential role in mitigating greenhouse gas
789 emissions, consistent with previous studies reporting vertical CO₂ flux balances observations in
790 semi-arid ecosystems (Rahimi et al., 2021; Tagesson et al., 2015; Agbohessou et al., 2023,
791 Wieckowski et al., 2024).

792 The estimated net C exchange budget-vertical balance is close to the reported mean for Sahelian
793 ecosystems (-1.6 ± 0.5 Mg C-CO₂ ha⁻¹; Tagesson et al., 2016). The EC-based net C exchange budget
794 balance (-1.8 ± 0.1701 Mg C-CO₂ ha⁻¹) is also similar to the value of -1.9 ± 0.4 Mg C-CO₂ ha⁻¹
795 reported for semi-arid savannas of northeastern Benin, despite higher annual rainfall (1495 mm;
796 Ago et al., 2016b). Furthermore, our EC estimate is close to the average net C exchange reported
797 for West African terrestrial ecosystems (-2.0 ± 1.5 Mg C-CO₂ ha⁻¹; Ago et al., 2016a).

798 However, estimates from Tagesson et al. (2015) (-2.7 ± 0.07 Mg C-CO₂ ha⁻¹) for a semi-arid
799 savannah in Dahra, Senegal, located between the 300 mm and 400 mm isohyets, were
800 comparatively higher. This is potentially attributable to specific characteristics of that specific
801 savannah site, such as herbaceous vegetation cover during the rainy season, the presence of
802 evergreen trees, and land management practices linked to pastoral livestock activities (Tagesson
803 et al., 2016).

804 The net annual C balance C exchange estimates presented in this study are, in fact, apparent fluxes
805 representing vertical fluxes only, given that they exclude organic matter (OM) imports and, more
806 critically, exports, introducing uncertainties. Notably, the export of crop residues and direct inputs
807 from animal excreta —particularly significant in 'bush fields' during the dry season — were not
808 accounted for. In our case of 'bush field', crop residues are exported to feed livestock, while

809 livestock faeces are collected for use as fuel or manure in 'home fields'. Such practices may lead to
810 a significant soil organic carbon stocks depletion (Malou et al, 2021), potentially diminishing the
811 net C budget (-1.4 ± 0.4602 Mg C-CO₂ ha⁻¹) over time and shifting the system closer to carbon
812 neutrality (Assouma et al., 2019).

813 These results should be contextualized within the broader framework of climate change and semi-
814 arid ecosystem management. Although agro-silvo-pastoral systems can function as **apparent**
815 annual carbon sinks, they remain highly sensitive to interannual rainfall variability and escalating
816 anthropogenic pressures. Sustainable management practices, particularly regarding **C**
817 **inputs/outputs from the system regarding crop harvest, residues export/exportations, and cattle**
818 **free manuring, must be taken into account are essential for to confirm-maintaining soil mineral**
819 **fertility and preserving the system's capacity of the system to act as effective a carbon sink,**
820 **thereby contributing to climate change mitigation.**

821 **4.7.4.8. Limitations of the study**

822 This study benefited from the inverse phenology of *F. albida*, allowing for direct comparison
823 between chamber-based GPP (GPPch.stand) and ecosystem-level GPP (GPP.EC) during the
824 leafless period of the trees. However, the system's spatial heterogeneity —common in
825 agroforestry— posed challenges for accurately partitioning CO₂ fluxes among trees, crops, and
826 soil. A key limitation was the development of weeds during the late rainy season, which
827 complicated the attribution of fluxes, particularly during the transitional period. Additionally,
828 while GPPch was successfully standardised by LAI for upscaling, this was not feasible for
829 respiration. Respiration integrates both autotrophic and heterotrophic components, which
830 respond to different drivers and are not directly linked to LAI, limiting the precision of upscaled
831 Rch.

832 Future improvements should aim to separately quantify respiration sources —tree roots, crops,
833 and microbial (heterotrophic) respiration— and account explicitly for the weed layer, to refine
834 flux partitioning in such complex agroforestry systems.

835 **Furthermore, the present study constitutes only an intermediate step delivering a first integrated**
836 **estimate of the main vertical CO₂ exchanges (photosynthesis, respiration, and net ecosystem**
837 **exchange) as a base for a forthcoming paper that will present a more comprehensive carbon**
838 **budget of the ecosystem. Establishing such a carbon budget would require substantial additional**
839 **data acquisition and poses considerable methodological challenges. In particular, quantifying**
840 **carbon inputs/outputs associated with free-ranging livestock grazing would be difficult to achieve**
841 **with acceptable accuracy. It must also be recognised that the system is in a dynamic, non-steady**
842 **state, characterised by marked inter-annual variability as well as periods of carbon storage and**
843 **release, which are difficult to constrain empirically except through modeling.**

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844 **Conclusion**

845 This study demonstrates the successful application of automated static chambers to quantify CO₂
846 fluxes in a Sahelian agroforestry system dominated by *F. albida*. The continuous, high-frequency
847 measurements captured key seasonal dynamics and short-lived events (e.g., Birch effect),
848 providing a more accurate assessment of carbon exchange than traditional intermittent sampling.
849 By integrating crop and soil components and applying dynamic partitioning models, the study
850 quantified both respiration and photosynthesis fluxes at fine temporal resolution. The results
851 revealed a clear 'fertile island' effect under tree canopies, with higher respiration and
852 photosynthetic activity, and highlighted the significant contribution of *F. albida* trees to annual
853 carbon uptake.

854 The consistency between chamber- and eddy covariance-based estimates reinforces the
855 robustness of the methodology. Overall, this work underscores the role of *F. albida*-based
856 agroforestry systems ~~in the dynamic of C exchanges as effective carbon sinks~~ in semi-arid
857 environments, offering valuable insights for carbon accounting and sustainable land management
858 in the Sahel.

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876 **Author contribution: CRediT**

877 **Seydina Mohamad BA:** Conducting in situ experiments, collecting and processing data,
878 writing-original draft, review and editing. **Olivier Roupsard:** Designing experimental
879 apparatus and methodology, writing, review and editing. **Lydie Chapuis-Lardy:** Designing
880 methodology, writing, review and editing. **Yélognissè Agbohessou:** Processing data,
881 review and editing. **Fred Bouvery:** Designing chambers and connection to the instrument,
882 review and editing. **Maxime Duthoit:** Designing experimental set and methodology,
883 review and editing. **Aleksander Wieckowski:** Review and editing. **Mohamed Habibou**
884 **Assouma:** Review and editing. **Espoir Gaglo:** Processing data, review and editing. **Claire**
885 **Delon:** review and editing. **Torbern Tagesson:** Designing methodology, review, and
886 editing. **Bienvenu Sambou:** Review and editing. **Dominique Serça:** Designing methodology,
887 writing, review and editing.

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