



Technical Note: Identifying Biomass Burning Emissions During ASIA-AQ Using Greenhouse Gas Enhancement Ratios

Jason A. Miech^{1,2}, Joshua P. DiGangi¹, Glenn S. Diskin¹, Yonghoon Choi^{1,3}, Richard H. Moore¹, Luke D. Ziemba¹, Francesca Gallo^{1,3}, Carolyn E. Jordan^{1,3}, Michael A. Shook¹, Elizabeth B. Wiggins¹, Edward L. Winstead^{1,3}, Sayantee Roy^{1,2}, Young Ro Lee⁴, Katherine Ball⁴, John D. Crounse⁴, Paul Wennberg⁴, Felix Piel⁵, Stefan Swift⁶, Wojciech Wojnowski^{5,7}, Armin Wisthaler^{5,6}

¹NASA Langley Research Center, Hampton, 23666, USA

²Oak Ridge Associated Universities, Oak Ridge, 37831-0117, USA

³Analytical Mechanics Associates, Hampton, 23666, USA

⁴California Institute of Technology, Pasadena, 91125, USA

⁵Department of Chemistry, University of Oslo, Oslo, 0313, Norway

⁶Institute for Ion Physics and Applied Physics, University of Innsbruck, Innsbruck, 6020, Austria

⁷Gdańsk University of Technology, Gdańsk, 80-233, Poland

Correspondence to: Jason A. Miech (jason.a.miech@nasa.gov)

Abstract. Biomass burning (BB) is a primary source of atmospheric chemistry reactants, aerosols, and greenhouse gases. Smoke plumes have air quality impacts local to the fire itself and regionally via long distance transport. Open burning of agriculture fields in Southeast Asia leads to frequent seasonal occurrences of regional BB-induced smoke haze and long-range transport of BB particles via the northeast monsoon. The Airborne and Satellite Investigation of Asian Air Quality (ASIA-AQ) campaign visited several areas including the Philippines, South Korea, Thailand, and Taiwan during a time of agricultural burning. This campaign consisted of airborne measurements on the NASA DC-8 aircraft aimed to validate observations from South Korea's Geostationary Environment Monitoring Spectrometer (GEMS) and to address local air quality challenges. We developed a method that used a combination of BB markers to identify ASIA-AQ DC-8 data influenced by BB and flag them. Specifically, we used rolling slope enhancement ratios of CO/CO₂ and CH₄/CO along with mixing ratios of CH₃CN, HCN, and CO, and particle scattering coefficient measurements. The flag was triggered when a combination of these variables exceeded a flight specific threshold. We found varying levels of BB-influence in the areas studied, with data flagged for BB being <1% for the Philippines and Korea, and <2% for Taiwan, but 19% for Thailand. Our method for flagging ASIA-AQ BB-affected data can be used to focus additional analyses of the ASIA-AQ campaign such as pairing with back-trajectories, satellite hotspot products, and microphysical aerosol characteristics.

1 Introduction

As Asian economies and populations continue to grow, so will their contribution to global greenhouse gas (GHG) emissions, driven primarily by increases in fossil fuel combustion and seasonally by biomass burning (BB) events. Left unchecked, these emissions will negatively impact air quality and climate, and therefore it is imperative that emission sources are properly identified and accounted for in emission inventories. Sources of carbon monoxide (CO) can include: domestic (residential



34 heating and cooking), industrial, and transport sectors; for methane (CH_4): fossil fuels (coal mining), agricultural emissions
35 (enteric fermentation and rice cultivation), and solid waste disposal and wastewater treatment; and for carbon dioxide (CO_2):
36 fossil fuel/biofuel combustion processes (powerplants and transportation) and industrial processes (cement production)
37 (Kurokawa et al., 2013; Kurokawa and Ohara, 2020). Biomass burning is also a significant source of these GHGs; however,
38 the seasonal nature of crop residue burning, and unpredictability of natural fires makes it difficult to accurately quantify these
39 variables sources (Akagi et al., 2011; Streets et al., 2003; Crutzen and Andreae, 1990).

40 The Airborne and Satellite Investigation of Asian Air Quality (ASIA-AQ) field campaign was conducted during February-
41 March 2024 near the maximum period of seasonal agricultural burning in Southeast Asia, providing an ideal opportunity to
42 further study biomass burning. ASIA-AQ was an international joint air quality campaign between the National Aeronautics
43 and Space Administration (NASA) and several space and environmental agencies in Korea (National Institute of
44 Environmental Research (NIER) and Korea Meteorological Administration (KMA)), the Philippines (Department of
45 Environment and Natural Resources (DENR), Philippine Space Agency (PhilSA), and Manila Observatory), Thailand (Geo-
46 Informatics and Space Technology Development Agency (GISTDA) and Pollution Control Department (PCD)), and Taiwan
47 (Ministry of Environment, National Central University (NCU), and Academia Sinica). Observations included in-situ
48 measurements of trace gasses, aerosol properties, radiation fluxes, and meteorological parameters from NASA's DC-8 aircraft
49 and several Korean aircraft, remote sensing measurements from NASA's G-III aircraft, satellite observations from Korea's
50 GEMS instrument, and several ground station measurements.

51 Current observation-based (top-down) techniques for GHG source apportionment include aircraft-based mass balances
52 (Cambaliza et al., 2015), standard eddy covariance (Helfter et al., 2016), positive matrix factorization (Guha et al., 2015), and
53 isotopic signatures (Fiehn et al., 2023). Enhancement ratios between GHGs have also been used to characterize sources of BB
54 plumes (Akagi et al., 2011; Andreae and Merlet, 2001; Nara et al., 2017; Yokelson et al., 1996); however, Yokelson et al.
55 (2013) pointed out several issues with this technique including problems that arise when background concentrations are
56 changing over the measurement period. Halliday et al. (2019) used short-term $\Delta\text{CO}/\Delta\text{CO}_2$ calculated over 60s rolling windows
57 and filtered data by the coefficient of determination (R^2) to negate the need for a consistent background measurement. DiGangi
58 et al. (2021) applied a similar method to data from NASA's Atmospheric Carbon Transport-America (ACT-America)
59 campaign to characterize local CO_2 emissions. (DiGangi et al., 2025) adapted this method for use with CH_4/CO ratios for
60 NASA's Cloud, Aerosol, and Monsoon Processes Philippines Experiment (CAMP²EX) airborne field campaign in a chemical
61 influence flag, with a focus on separating BB and urban influences. Previous work during the Fire Influence on Regional to
62 Global Environments and Air Quality (FIREX-AQ) airborne campaign used CO and black carbon (BC) enhancements to
63 identify and flag smoke plumes (Warneke et al., 2023). Since the ASIA-AQ campaign overflew many urban areas, a more BB-
64 specific approach was required. The method presented in this work combines the use of CO/CO_2 and CH_4/CO ratios to further
65 constrain source contributions. To specifically target BB sources, we have also incorporated mixing ratios of acetonitrile
66 (CH_3CN) (Yokelson et al., 2009), hydrogen cyanide (HCN) (Yokelson et al., 2009), CO mixing ratio (Lin et al., 2013; Nara et
67 al., 2017), and aerosol scattering coefficient (Lin et al., 2013; Tsay et al., 2013) into our method. This work provides a targeted



method to identify air masses influenced by biomass burning during the 2024 ASIA-AQ airborne field campaign, an example of which is shown in our Thailand case study.

2 Data and Methods

2.1 ASIA-AQ Field Campaign and In Situ Aircraft Measurements

These methods were developed for the ASIA-AQ campaign, which was a joint international field study focused on air quality challenges local to its areas of study, which included the Philippines, Korea, Thailand, and Taiwan. Data for this campaign were collected from aircraft, including NASA's DC-8 and G-III and three Korean aircraft, and several ground sites. Sixteen DC-8 science flights were conducted between 6 February 2024 and 27 March 2024; the full flight break-down can be found in Table S1 along with typical flight paths in Figures S1-4. The NASA DC-8 aircraft carried a variety of instruments used for in-situ gas-phase, aerosol, radiation, and meteorological measurements. CO₂ was measured via non-dispersive infrared spectroscopy using a modified LICOR 7000 instrument with an uncertainty of 0.25 ppm at <500 ppm and 2% at > 500 ppm (Vay et al., 2003). CH₄ and CO were measured via wavelength modulation spectroscopy using the Differential Absorption Carbon monoxide Measurement (DACOM) instrument, with CO uncertainties of 2% for < 1 ppm and 5% for > 1 ppm and CH₄ uncertainty of 1% (Sachse et al., 1987, 1991). HCN mixing ratios were measured using a Chemical Ionization High Resolution Time of Flight Mass Spectrometer with an uncertainty of 25% plus 35 pptv (CIT-CIMS) (Crounse et al., 2006). CH₃CN measurements were taken using a Proton-Transfer-Reaction Time-of-Flight Mass Spectrometer (PTR-MS) with an uncertainty of 13% (Müller et al., 2016; Reinecke et al., 2023). Total aerosol scattering at 550 nm measurements were measured using a TSI-3563 Nephelometer with an estimated accuracy of 20% and estimated precision of 1 Mm⁻¹ (Anderson and Ogren, 1998). Accumulation-mode aerosol BC mass concentrations were measured by a Single Particle Soot Photometer (SP2-D) (Droplet Measurement Technologies Inc. (DMT), Longmont, Colorado, USA) with an uncertainty of 20%, operated by NASA Langley Research Center. Optical particle size distributions for particles with diameters between 63.1 and 1000 nm were measured using a DMT Ultra High Sensitivity Aerosol Spectrometer (UHSAS) Model UHSAS-0.055 (DMT, Longmont Colorado USA) with an uncertainty of 20%, while particles with diameters between 3.16 and 89.1 nm were sized using a Scanning Mobility Particle Sizer (SMPS) with an uncertainty of 20%. Submicron particle number concentrations (>10 nm) were measured with a TSI Condensation Particle Counter (CPC) CPC-3772 (TSI Inc., Shoreview, Minnesota, USA) with an uncertainty of 10%. A duplicate instrument also with 10% uncertainty was used to measure non-volatile particle number concentrations (>10 nm) with the sample passing through a thermal denuder heated to 350 °C prior to entering the instrument. All DC-8 data used in this work had a time resolution of 1 Hz, with the exceptions of the SMPS which measured every minute. Hourly Thai precipitation measurements were obtained from the Thailand Pollution Control Department using the Phaya Thai, Khet Phaya Thai, and Bangkok sites. This data and the developed BB flag are publicly available on the ASIA-AQ data archive (ASIA-AQ Science Team, 2024).



99 2.2 Rolling Slope Method

100 The rolling slope calculation method employed here was taken from Halliday et al. (2019), where linear fits over 120s rolling
101 windows of CO vs. CO₂ and CH₄ vs. CO were calculated using error adjusted bivariate regression as detailed in York et al.
102 (2004) and Cantrell (2008). Applying this method, a maximum of 121 observations were used in each calculation, while the
103 minimum for a valid calculation was set at three. Slopes with low goodness of fit ($R^2 < 0.5$) were filtered out as uncorrelated
104 and, slopes with a ΔCH_4 , ΔCO , or ΔCO_2 less than five times the precision value were dropped. While Halliday et al. (2019)
105 demonstrated that the shapes of the slope distributions are somewhat insensitive to the R^2 cutoff (when $R^2 \geq 0.5$), here we
106 observed differences between correlated slopes ($R^2 \geq 0.5$) and uncorrelated slopes ($R^2 < 0.5$). Halliday et al. (2019) also showed
107 that varying the window width did not drastically change the slope distributions. Similarly, Figures S5-7 show that with
108 increasing window width, the number of correlated slopes increases and with an increasing R^2 cutoff the number decreases;
109 however, the shapes of the slope distributions are similar. Table S2 exhibits the percentage of calculated $\Delta\text{CO}/\Delta\text{CO}_2$ and
110 $\Delta\text{CH}_4/\Delta\text{CO}$ slopes that were correlated for each flight. On average across all flights, 52% of $\Delta\text{CO}/\Delta\text{CO}_2$ slopes and 53% of
111 $\Delta\text{CH}_4/\Delta\text{CO}$ slopes were correlated according to the $R^2 \geq 0.5$ criterion, in 120s rolling windows.

112
113 $\Delta\text{CO}/\Delta\text{CO}_2$ rolling slopes can be used to isolate plumes or air mass boundaries with correlated behavior and provide a
114 rudimentary source classification. The combustion efficiency of the source is inversely related to the $\Delta\text{CO}/\Delta\text{CO}_2$ slope,
115 therefore we have divided up source behaviors into four combustion efficiency bins based on $\Delta\text{CO}/\Delta\text{CO}_2$ slopes: 0-1%, 1-2%,
116 2-4%, and >4%, as previously used by Halliday et al., 2019. In this work, the >4% combustion efficiency bin is most relevant
117 for identifying BB-influenced plumes. Figure 1A shows an example of this classification applied to the CO and CO₂ bulk
118 ratios from Thailand Flight 1 demonstrating the ability of this method to highlight different air mass behaviors. $\Delta\text{CH}_4/\Delta\text{CO}$
119 slopes have been used to distinguish between urban sources of methane and biomass burning (DiGangi et al., 2025; Reid et
120 al., 2023), with higher slopes indicative of more urban sources and smaller slopes $0 < x < \sim 40\%$ with biomass burning. Figure
121 1B displays the application of these regimes to CO and CH₄ bulk ratios. Both Fig. 1A and 1B demonstrate a large influence
122 from slopes indicative of biomass burning, >4% for $\Delta\text{CO}/\Delta\text{CO}_2$ (shown in dark red) and $0 < x < \sim 40\%$ for $\Delta\text{CH}_4/\Delta\text{CO}$ (shown
123 in orange).

124

125

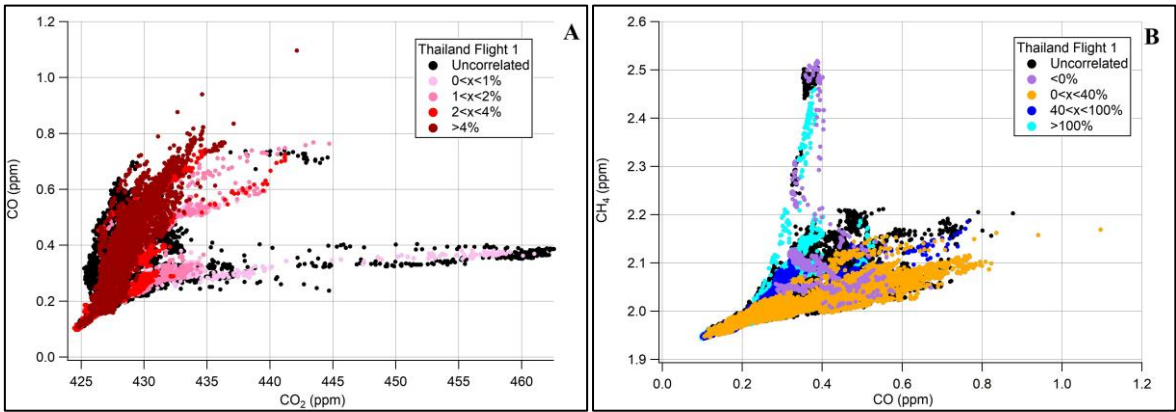


Figure 1. CO vs CO₂ mixing ratios colored by $\Delta\text{CO}/\Delta\text{CO}_2$ rolling slope enhancement ratio regimes (A) and CH₄ vs CO mixing ratios colored by $\Delta\text{CH}_4/\Delta\text{CO}$ rolling slope enhancement ratio regimes (B) for Thailand flight 1.

2.3 Biomass Burning Flag Determination

The biomass burning flag uses a combination of variables indicative of or a product of biomass burning, including CO mixing ratio, particle scattering at 550 nm, HCN, CH₃CN, $\Delta\text{CO}/\Delta\text{CO}_2$, and $\Delta\text{CH}_4/\Delta\text{CO}$. The BB flag is triggered when at least two of the variables meet their flight-specific thresholds (Table S3). However, when CO mixing ratio and particle scattering are paired together a third variable needs to meet its threshold to trigger the flag; the reasoning for this is to prevent false positives occurring from inefficient combustion processes other than biomass burning. Additionally, the BB flag is triggered when all four non-slope variables are within 10% of their threshold to address edge cases. Example scenarios for triggering the flag are shown in Table 1.

Table 1. Example Scenarios for triggering the BB flag for Thailand flight 1 (20240316). Underlined values in green represent the variables that have exceeded their threshold, italic values in blue are variables within 10% of their threshold. The last row represents an example where all four non-slope variables are at least within 10% of their threshold.

BB Flag	$\Delta\text{CH}_4/\Delta\text{CO}$ (%)	$\Delta\text{CO}/\Delta\text{CO}_2$ (%)	CO (ppm)	HCN (ppt)	CH ₃ CN (ppb)	Total scattering @ 550 nm (Mm ⁻¹)
Threshold	44.5	4	0.32	2750	1.50	160
✓	N/A	<u>5.1</u>	<u>0.67</u>	<u>2810</u>	1.31	<u>487</u>
✓	<u>40.3</u>	<u>9.8</u>	<u>0.34</u>	1310	0.56	<u>227</u>
✗	54.6	2.5	0.22	774	0.30	36
✓	N/A	N/A	<u>0.53</u>	<u>2823</u>	0.91	<u>386</u>
✗	N/A	N/A	<u>0.33</u>	1053	0.57	<u>186</u>
✓	N/A	2.2	<u>0.67</u>	<u>2715</u>	<u>1.43</u>	<u>418</u>



2.3.1 Variable Threshold Determination

The BB flag shares similarities with the Chemical Influence flag developed for CAMP²EX, where lower positive $\Delta\text{CH}_4/\Delta\text{CO}$ rolling slopes were associated with biomass burning influence (DiGangi et al., 2025; Reid et al., 2023). Here, the specific $\Delta\text{CH}_4/\Delta\text{CO}$ cutoff values were chosen on a flight-by-flight basis by evaluating the slope distribution for each flight and looking for distinctive populations. An example of this process is shown in Fig. 2A, where there is a distribution separation occurring at a slope value of 44.5% for the first Thailand flight. For $\Delta\text{CO}/\Delta\text{CO}_2$ the threshold was set to greater than 4% for all flights, as low efficiency processes like biomass burning are associated with higher $\Delta\text{CO}/\Delta\text{CO}_2$ values (Halliday et al., 2019). The thresholds for the other variables were initially determined by identifying at which concentration within the biomass burning regime (as set by $\Delta\text{CH}_4/\Delta\text{CO}$) the variable is distinguishable from non-biomass burning influence regimes. In the Fig. 2B example, that regime is from 0-44.5% $\Delta\text{CH}_4/\Delta\text{CO}$, as determined from Fig. 2A, and the CH_3CN threshold was set to 1.00 ppb. This method was repeated for the other variables (HCN, particle scattering, and CO mixing ratio) and for all flights. Thresholds for the mixing ratio and scattering variables were set on a flight-by-flight basis.

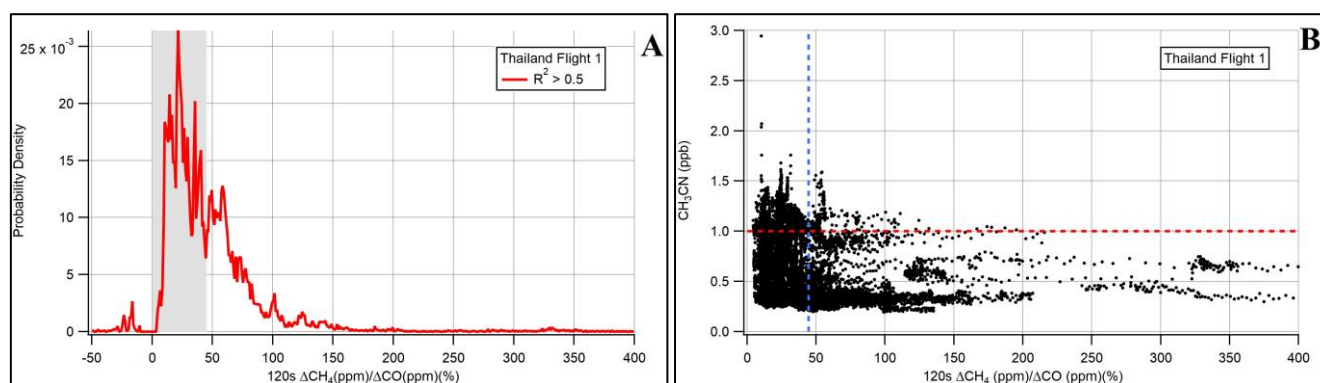


Figure 2 (A) Frequency distribution of rolling slope $\Delta\text{CH}_4/\Delta\text{CO}$ enhancement ratios. The shaded region is the set BB regime for this flight, as determined by the minimum in the distribution. (B) CH_3CN vs $\Delta\text{CH}_4/\Delta\text{CO}$ for Thailand flight 1. The blue vertical line represents the edge of the biomass burning influence regime at 44.5% $\Delta\text{CH}_4/\Delta\text{CO}$, while the red horizontal line is the preliminary CH_3CN threshold for the BB flag.

In the case of thresholds that were more difficult to distinguish, a higher value was initially selected to minimize false positives in the flag. Thresholds were then further refined from this initial value by examining the sensitivity of the total percentage of points flagged to that threshold. The final thresholds were chosen at the lowest levels for which the percentage of points remained essentially constant. Figure 3 demonstrates the results from one of these sensitivity tests performed on the first flight in Thailand. In this case, decreasing the thresholds for CO mixing ratio, CH_3CN , HCN, and particle scattering results in large increases in the number of points flagged. In Fig. 3A, the dashed HCN and CH_3CN traces are still decreasing past their initial threshold values represented by the triangles (2400 ppt and 1.0 ppb); therefore, this is motivation to increase their thresholds by ~15% and ~50% respectively, resulting in the solid traces. The initial thresholds for CO mixing ratio and particle scattering (0.32 ppm and 160), shown as triangles Fig. 3B required no adjustment for this particular flight; the decrease in the percentage



of points flagged for the post-adjustment traces is due to the increases in the HCN and CH₃CN thresholds. The final determined thresholds and rolling slope regimes for all flights are in Table S3, while Table 1 provides some example scenarios of the BB flag being triggered or not for Thailand flight 1. The concentration thresholds varied by a factor of 2.6 for CO and up to 8 for CH₃CN, this variability can be explained by changing background concentrations for each flight day and location. This approach was followed to minimize false positives in the dataset; therefore, this method is less sensitive to air masses with only minor influences from biomass burning.

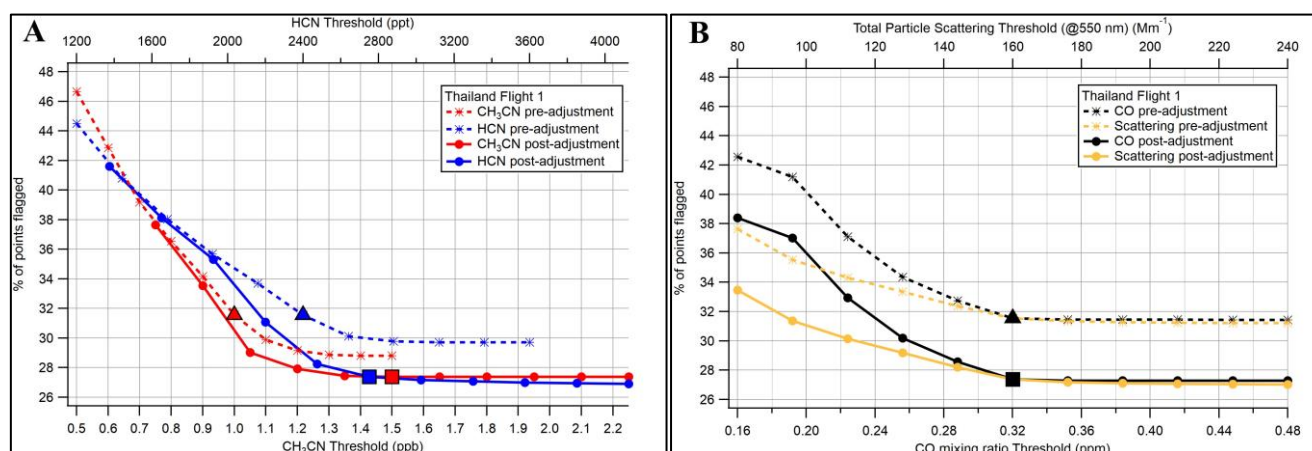


Figure 3. Example variable threshold sensitivity plots for the first Thailand flight. Fig 3A shows the sensitivities for the HCN and CH₃CN thresholds established using plots like Fig 2B in the dashed traces and the final thresholds in solid. Fig 3B shows the sensitivities for the CO mixing ratio and total particle scattering thresholds established using plots like Fig 2B in the dashed traces and the final thresholds in solid. The triangles represent the initial thresholds for each variable and the squares the final thresholds.

2.4 HYSPLIT Back-Trajectories

Air mass history was probed using 48 hour back-trajectories calculated from the DC-8 flight track at one second intervals using NOAA's HYSPLIT model using GFS 0.25° meteorology (Draxler, 1999; Draxler and Hess, 1997, 1998; Stein et al., 2015). For more specific source type, altitude, and receptor location analysis, we used the BB flag as a filter and focused on specific areas and altitudes of the flight path, such as urban areas at low altitudes to determine whether the emissions measured in these locations were local or transport. To empirically quantify these contributions, we totaled the number of back-trajectory points under 1 km that traveled through a certain area prior to ending up at the receptor location; locations used in this analysis are described in Table S4 and shown in Figure S8.

2.5 VIIRS Fire Hotspots & Imagery

The Visible Infrared Imaging Radiometer Suite (VIIRS) I-Band 375 m Active Fire data product was used to assess fire hotspot density over Southeast Asia (Schroeder et al., 2014). Data version 2.0 from the Suomi National Polar-Orbiting Partnership (Suomi NPP) spacecraft were downloaded from NASA-FIRMS (Fire Information for Resource Management System) (NASA FIRMS, 2024). VIIRS Corrected Reflectance (True Color) imagery was downloaded from NASA-FIRMS (Lin and Wolfe,

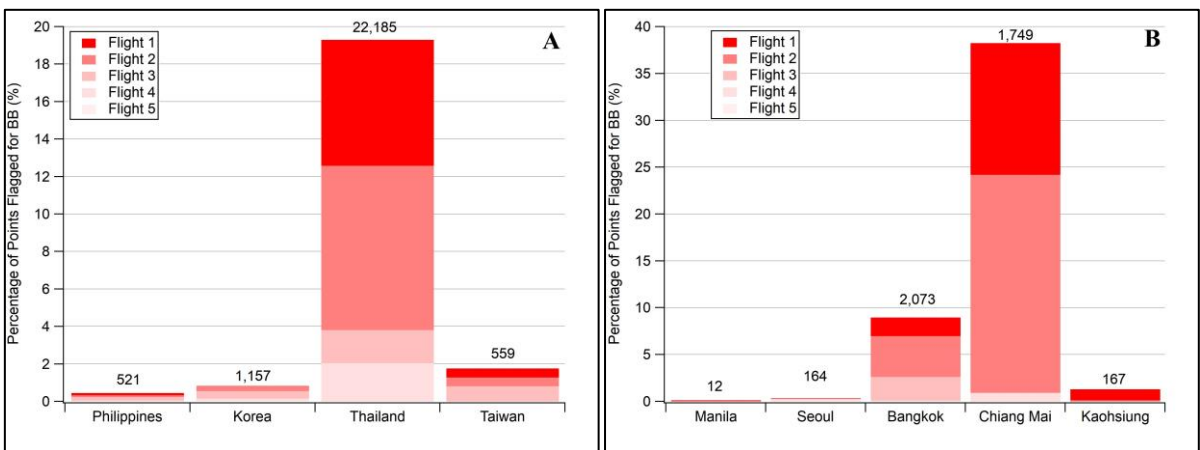


2022a, b; NASA VIIRS Characterization Support Team (VCST)/ MODIS Adaptive Processing System (MODAPS), 2022a, b). The Ozone Mapping and Profiling Suite (OMPS) Aerosol Index overlay was downloaded from NASA-FIRMS (Torres and Goddard Earth Sciences Data and Information Services Center (GES DISC), 2019).

3 Results and Discussion

3.1 Breakdown of points flagged for each area

Figure 4A breaks down the percentage of points flagged for BB for each flight and for each location, with a breakdown based on cities in Fig 4B. Very little BB-influence was observed in the Philippines, with most of it occurring north of Manila under 1.5 km above ground level (AGL) (Fig S1). For Korea, the highest occurrences of BB occurred on flights 2 and 3. On flight 2 the majority of the BB was observed over the West Sea under 500 m AGL. For flight 3 there was BB-influence over both the West Sea and Seoul above 1 km AGL (Fig S2). Most of the observed BB-influence for the campaign occurred in Thailand, specifically flights 1 and 2, with Chiang Mai experiencing more BB compared to Bangkok (Fig S3), except for flight 3. Section 3.2 delves more into the geographical and temporal BB flag differences observed in Thailand during the campaign. Taiwan had the next highest percent occurrence, but Korea overall had more points flagged for BB compared to Taiwan. The total time spent collecting data over these locations is the reason for this discrepancy, the campaign spent almost four times as much time over Korea compared to Taiwan. However, in terms of cities, Kaohsiung had the third highest percentage and number of points flagged, after Chiang Mai and Bangkok, with most of these flagged points occurring on flight 1. Most of the BB observed in Taiwan occurred during the third flight, with the majority occurring on parts of the flight track not over Kaohsiung and above 1 km AGL (Fig S4).



211
212 **Figure 4. Percentage of points flagged for BB for each flight over each country (A) and select cities (B) with total number of points**
213 **flagged displayed.**



214 3.2 Thailand Case-Study

215 3.2.1 Flight by Flight Breakdown

216 Since Thailand had the largest prevalence of biomass burning among the flight locations, we have broken down these flights
217 in greater detail. Figure 5 shows the flight-by-flight breakdown of the BB flag (A-D), VIIRS 48h fire hotspot density maps (E-
218 H) and VIIRS satellite imagery with the OMPS aerosol index overlay (I-L) for the four Thailand flights. While flight 2 had
219 the highest overall occurrence of BB (Fig 5B), flight 1 had more BB-influence closer to the surface (Fig 5A). An increase in
220 fire hotspot density is also observed when comparing Fig 5E to 5F, where the density has increased in Eastern Thailand,
221 Cambodia, and Northern Thailand/Myanmar. However, a similar increase is not observed looking at the OMPS aerosol index
222 overlays (Fig 5I to Fig 5J). For flight 1 (Fig 5I) the aerosol index is elevated over northern Laos, northern Vietnam, northern
223 and western Thailand, and central and eastern Myanmar, indicating either a more concentrated, thicker and/or higher layer of
224 absorbing aerosols (dust and smoke) in these areas. For flight 2 (Fig 5J) the higher aerosol index values are concentrated over
225 Hanoi, with western Thailand clearing up and few elevated values over eastern Thailand, Cambodia, and southern Laos. This
226 apparent aerosol index decrease around Thailand contrasts with what was observed using the BB flag and the VIIRS fire
227 hotspot density maps, however, the OMPS aerosol index is also sensitive to changes in aerosol height, concentration, and
228 degree of absorption (Torres and Goddard Earth Sciences Data and Information Services Center (GES DISC), 2019). Prior to
229 flight 3, the area received a considerable amount of rain (~41 mm in Bangkok and up to 82 mm on the northern flight track on
230 the day before flight 3), which likely washed out most of the smoke and put out local fires, seen in the decreased fire hotspot
231 density across Thailand (Fig 5G), and decreased aerosol indices (Fig 5K). For the final flight, there's a slight increase in the
232 number of points flagged compared to flight 3 (~17%), more near the surface as burning resumed, which is apparent when
233 comparing the VIIRS fire hotspot density maps (Fig 5G and Fig 5H) and scattered elevated aerosol indices over northern
234 Thailand (Fig 5L).

235 Figure 5A-D also shows a breakdown of BB-influence in terms of major population centers on the flight track, specifically
236 Bangkok on the southern portion of the track and Chiang Mai on the northern part. For flight 1, we observed no BB-influence
237 below 1 km in Bangkok across four low passes. This contrasts with observations during one pass in Chiang Mai where ~80%
238 of the low altitude data was flagged for BB. This is consistent with Fig 5E, where the fire hot spot densities are higher around
239 Chiang Mai compared to Bangkok, and Fig 5I with elevated aerosol indices over Chiang Mai. For flight 2, there was equivalent
240 amounts of BB-influence over the two cities but significantly more at the surface in Chiang Mai, however, relative to the time
241 spent at each location over 90% of data from Chiang Mai were flagged while only ~16% of Bangkok data were. Even though
242 the fire hotspot density around Bangkok is higher for flight 2 (Fig 5F) compared to flight 1 (Fig 5E), that BB-influence does
243 not appear to be reaching the surface. One possible explanation is that the prevailing southerly winds are blowing that smoke
244 more towards Central Thailand. In the third flight, the rain likely washed out almost all of the smoke around Chiang Mai and
245 quelled many fires as observed in Fig 5G. However, there was still some BB-influence observed above Bangkok, perhaps by
246 a transported airmass aloft unaffected by the rain. From Fig 5G, there are still high fire hotspot densities occurring in Cambodia



247 and Laos, which could be traced forward to Bangkok using Hysplit back-trajectories (Fig 6A). For the last flight, there was
248 little BB-influence observed in both cities ($<4\%$) but still a meaningful amount across the rest of the flight track ($\sim 8\%$). Figure
249 5H shows that fire activities seemed to have resumed throughout Thailand in the 48 hours leading up to flight 4.

250
251
252
253
254
255
256
257
258
259
260
261
262
263
264
265
266
267
268
269
270
271
272
273
274

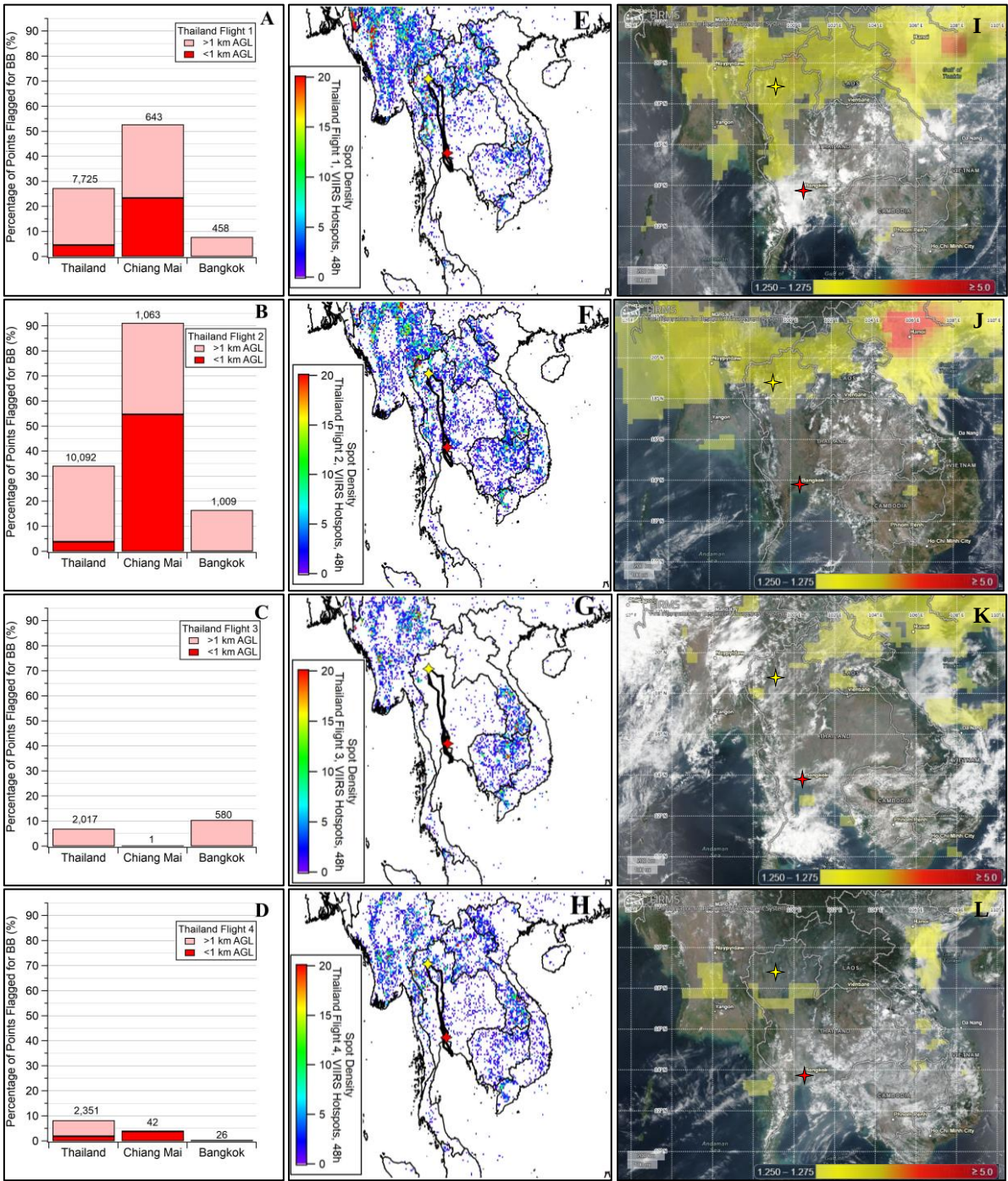


Figure 5. (A-D) Percentage of points flagged for BB for each Thailand flight across the whole flight track, over Chiang Mai, and over Bangkok for above and below 1 km AGL with total number of points flagged displayed. (E-H) 48h VIIRS fire hotspot density maps over Southeast Asia for March 16, 18, 21, and 25 2024. The DC-8 flight track is shown in black. The yellow star represents Chiang Mai, and the red star represents Bangkok. (I-L) VIIRS Corrected Reflectance (True Color) imagery with OMPS Aerosol Index overlay over Southeast Asia for March 16, 18, 21, and 25 2024.



284 3.2.3 Air Mass Origin Breakdown

285 To better understand the air mass history at these locations, we used Hysplit back-trajectory modeling along the flight path to
286 determine where these air masses came from and traveled through (using criteria from Table S4). Back-trajectories with an
287 altitude < 1 km were then quantified over certain areas of interest for portions of the flight track that passed over Chiang Mai
288 and Bangkok and were flagged for BB. Figure 6 shows the results for Bangkok for all of the flights above and below 1 km of
289 altitude. For the first two Bangkok flights, there was influence from the south (Gulf of Thailand and Malay Peninsula) and
290 northeast (Eastern Thailand) above 1 km altitude, which switched to the east (Cambodia and Vietnam) for flight 3. Under 1
291 km altitude for Bangkok flights 2 and 4, most of the BB was local or traveled from the south, while for flight 3 BB was
292 transported from the northeast (East Thailand and South China Sea). However, it should be noted that Bangkok experienced
293 <0.5% of BB-flagged data for Flight 4 (Fig 5D) so these results are mostly likely reflective of a single plume. Fig. 7 shows the
294 Hysplit back-trajectory quantifications for Chiang Mai. For Chiang Mai above 1 km of altitude, the majority of back-
295 trajectories are either from Central Thailand or Myanmar, with flight 1 dominated by BB transported from Central Thailand
296 and flight 2 by Myanmar. Under 1 km altitude, there is a similar trend with flight 1 flagged back-trajectories originating from
297 south of Chiang Mai and flight 2 and 4 from the west. These results are in agreement with satellite fire hotspot retrievals (Fig.
298 5E-H), that show the highest hotspot density around and northwest of Chiang Mai for flights 1, 2, and 4 but for flight 3 (Fig
299 5G) a clearer area around Chiang Mai. Additional areas of high fire hotspot density occurred in Cambodia, Vietnam, and Laos,
300 especially in the last three flights.

301
302
303
304
305
306
307
308
309
310

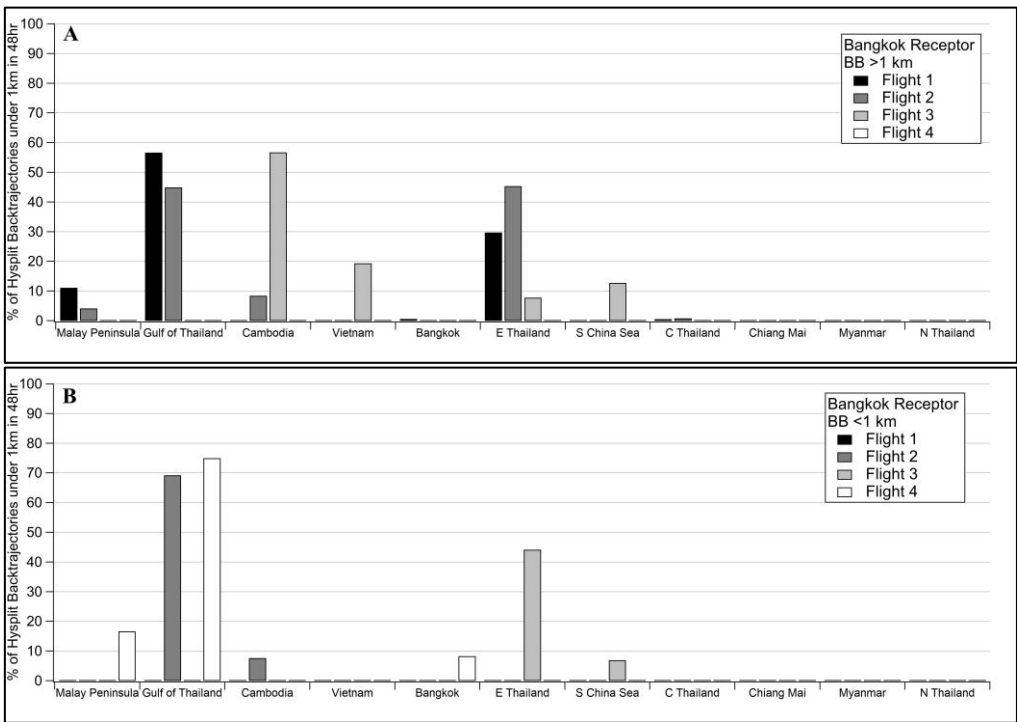


Figure 6. Hysplit back-trajectory quantification for BB-flagged data over Bangkok for all flights above (A) and below (B) 1 km AGL.

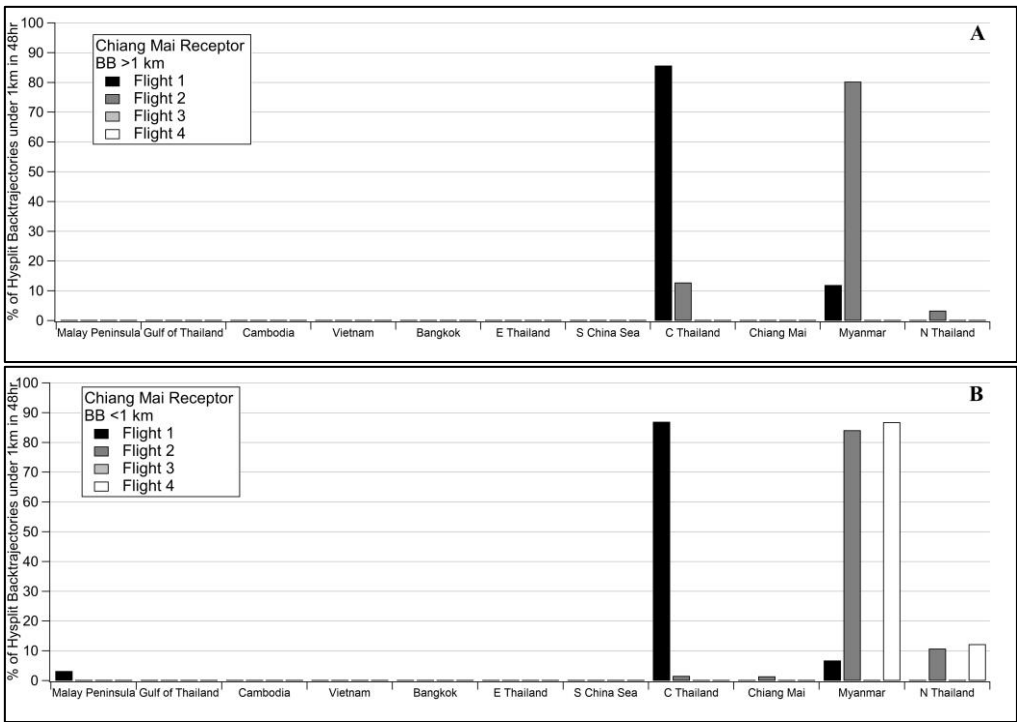


Figure 7. Hysplit back-trajectory quantification for BB-flagged data over Chiang Mai for all flights above (A) and below (B) 1 km AGL.



3.2.4 Comparison of Microphysical Aerosol Characteristics

To further illustrate the utility of the BB flag we have used it to compare the integrated size distribution of submicron aerosol particles. Figure 8A shows particle size distributions for BB-flagged data under 1 km in Chiang Mai and unflagged data under 1 km in Bangkok for the same day (flight 2). The BB-flagged data has a unimodal size distribution with a mean particle diameter of 150 nm (accumulation mode) compared to a bimodal size distribution in the unflagged data with a mean particle diameter 27 nm (nuclei mode). The second and less significant peak in the unflagged accumulation mode (mean diameter 136 nm) data may be the result of background BB presence, demonstrating that this flag is optimized for clear biomass burning influence and may overlook less obvious BB-influence. Several studies have demonstrated that aerosol size distributions in fresh (<1 h) smoke will start at median diameters ranging from 40-150 nm and grow to larger sizes with a decrease in modal width as they age (Hodshire et al., 2021; Janhäll et al., 2010; Reid et al., 1998). Figure 8B demonstrates a comparison between the same location (Chiang Mai <1 km) but on flights with differing amounts of BB-influence, i.e. flights 2 and 3. While the flight 3 unflagged data still exhibits smaller particle diameters (mean particle diameter 41 nm) than the flight 2 BB-flagged data (mean particle diameter 150 nm), it demonstrates a relatively stronger bimodal distribution compared to flight 2 unflagged data, with a mean particle diameter of 132 nm in the accumulation mode. Even though the presence of this accumulation mode peak and BB markers (CO, HCN, and CH₃CN) within 20% of their respective thresholds indicates an influence from BB, the dominant nuclei mode peak and $\Delta\text{CO}/\Delta\text{CO}_2$ (0.38%) and $\Delta\text{CH}_4/\Delta\text{CO}$ (104%) slopes point to urban combustion sources as the dominant influence for this unflagged data.

When looking at black carbon mass concentrations, shown in Fig 9A, the BB-flagged data has consistently higher BC concentrations compared to unflagged data across Thailand for flights 1 and 2. Flights 3 and 4 had the lowest BB-flagged BC concentrations and were higher compared to the unflagged data but were within one standard deviation. This is consistent with Fig 5C and 5D where flights 3 and 4 had the smallest number of points flagged for BB. The unflagged BC concentrations for flights 1, 2, and 4 are also higher than unflagged flight 3 (but within one standard deviation); therefore, assuming other BC emissions are consistent between the flight days, this is further evidence that the BB flag does not capture all BB-influenced data, specifically BB-influence in the background. Figure 9B shows that the non-volatile number fraction of fine particles for BB-flagged data is enhanced, with a narrower range compared to unflagged data across all four flights. One explanation may be that the increased BC mass concentration is elevating the non-volatile number concentration of BB-flagged data in combination with a decrease in the number concentration of volatile fine particles due to evaporation during transport. Future work examining the age of the smoke using short-lived BB tracers could provide more information on the concentration of volatile fine particles in fresh and aged smoke plumes.

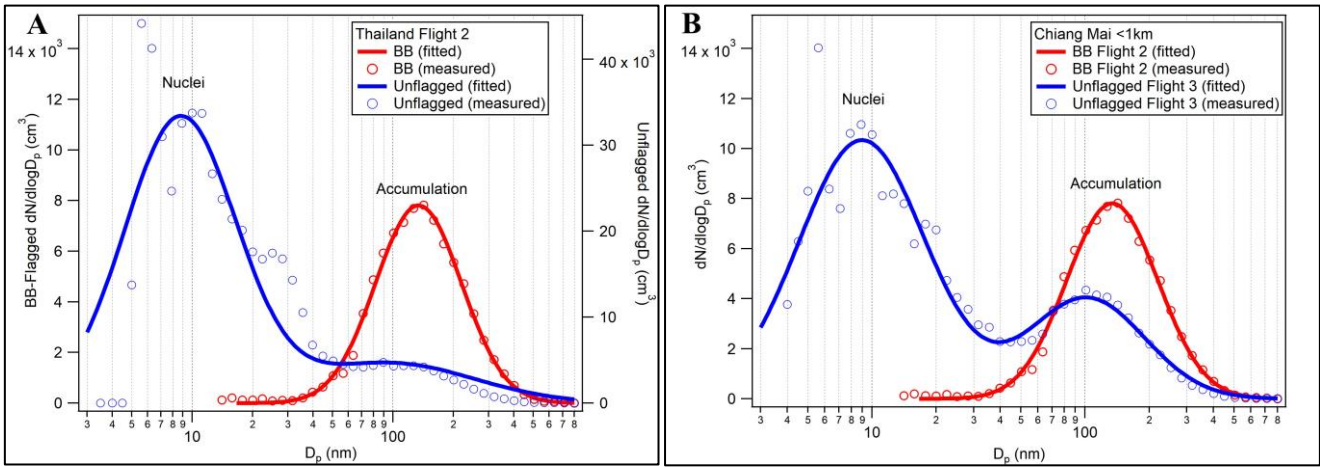


Figure 8. Fitted and measured particle size distributions for BB-flagged and unflagged conditions for flight 2 (A) and for Chiang Mai (B) < 1 km AGL. It should also be noted that the BB-flagged and unflagged distributions in 8A are on different scales but the same scale in 8B.

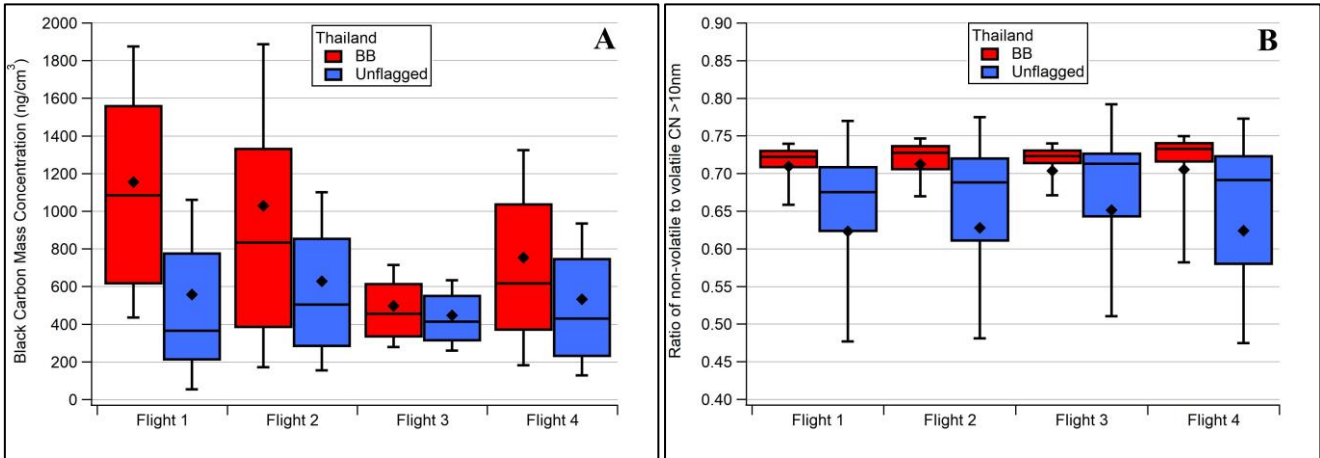


Figure 9. Black carbon mass concentrations (A) and non-volatile number fraction CN > 10 nm (B) for BB-flagged and unflagged data for each Thailand flight. Whiskers are representative of one standard deviation, while the boxes represent the interquartile range, the horizontal line the median, and the diamond the mean.

4 Conclusions

Biomass burning and inefficient combustion contribute to poor air quality and greenhouse gas emissions across the globe. Airborne investigations in regions prone to these emissions provide more detailed and focused measurements than typical ground or satellite methods. This work demonstrated a novel approach for identifying biomass burning-impacted airmasses in an airborne dataset. A combination of biomass burning tracers and indicators were used to distinguish biomass burning-impacted airmasses along each given flight track. The Thailand case study demonstrates the efficacy of this flag in determining areas most influenced by biomass burning and in combination with trajectory models, an idea of air mass history. Additionally,



a preliminary analysis of physical characteristics of aerosols revealed differences between BB-flagged and unflagged airmasses, including aerosol size distributions, black carbon concentrations, and non-volatile number fraction. These findings, while applicable to Asia, demonstrate the value of the method, which can be applied to other field campaigns with similar measurements. The utility of the BB flag can be increased in the future with the use of specific volatile organic compounds (VOCs) to provide information on the age of the smoke, e.g., primary BB VOCs versus secondary BB VOCs (Liang et al., 2022). Additional development of a boundary layer flag for the ASIA-AQ dataset will improve future work studying the health impacts of biomass burning smoke and inefficient combustion at ground level and transport in aloft layers.

Data Availability

All data used in this publication are open access and can be found in the ASIA-AQ Data Archive (<https://doi.org/10.5067/SUBORBITAL/ASIA-AQ/DATA001>), the FIRMS Archive (<https://firms.modaps.eosdis.nasa.gov/download/>), and the HYSPLIT model is available at <https://www.ready.noaa.gov/HYSPLIT.php>.

Author Contribution

JAM, JPD, GSD, YC, RHM, LDZ, FG, CEJ, MS, EBW, ELW, SR, YRL, KB, JDC, PW, FP, SS, WW, and AW participated in the data collection. JAM, JPD, GSD, MS, FG, KB, and WW performed the data analysis and provided feedback. JAM prepared the manuscript with contributions from all coauthors.

Competing Interests

The contact author has declared that none of the authors have any competing interests.

Acknowledgments

The authors would like to thank Jim Crawford and the entire ASIA-AQ leadership team for their support, guidance, and feedback. The authors thank NASA's Earth Science Project Office for their logistical support collecting this data. We thank Ryan Bennett, David Van Gilst, and Terry Hu for aircraft navigational and meteorological data. The authors thank Dave Eckberg and Mauro Rana for their software and hardware support with the DACOM-DLH instrumentation. Jason Miech's research was supported by an appointment to the NASA Postdoctoral Program at the NASA Langley Research Center, administered by Oak Ridge Associated Universities under contract with NASA. Data collection was funded by the ASIA-AQ project under NASA's Tropospheric Composition Program. PTR-ToF-MS measurements aboard the NASA DC-8 during ASIA-AQ were partially funded by the Austrian Federal Ministry for Climate Action, Environment, Energy, Mobility, Innovation, and Technology (BMK), represented by the Austrian Research Promotion Agency (FFG), through the Austrian Space Applications Programme (ASAP 2022, #FO999900547). IONICON Analytik is acknowledged for supplying a FUSION PTR-ToF-MS analyzer and providing staff support. The authors gratefully acknowledge the NOAA Air Resources Laboratory



(ARL) for the provision of the HYSPLIT transport and dispersion model and/or READY website (<https://www.ready.noaa.gov>) used in this publication. We acknowledge the use of data and/or imagery from NASA's Land, Atmosphere Near real-time Capability for Earth observations (LANCE) (<https://earthdata.nasa.gov/lance>), part of NASA's Earth Science Data and Information System (ESDIS). We acknowledge the use of data and/or imagery from NASA's Fire Information for Resource Management System (FIRMS) (<https://www.earthdata.nasa.gov/data/tools/firms>), part of NASA's Earth Science Data and Information System (ESDIS).

References

- Akagi, S. K., Yokelson, R. J., Wiedinmyer, C., Alvarado, M. J., Reid, J. S., Karl, T., Crounse, J. D., and Wennberg, P. O.: Emission factors for open and domestic biomass burning for use in atmospheric models, *Atmospheric Chemistry and Physics*, 11, 4039–4072, <https://doi.org/10.5194/acp-11-4039-2011>, 2011.
- Anderson, T. L. and Ogren, J. A.: Determining Aerosol Radiative Properties Using the TSI 3563 Integrating Nephelometer, *Aerosol Science and Technology*, 29, 57–69, <https://doi.org/10.1080/02786829808965551>, 1998.
- Andreae, M. O. and Merlet, P.: Emission of trace gases and aerosols from biomass burning, *Global Biogeochemical Cycles*, 15, 955–966, <https://doi.org/10.1029/2000GB001382>, 2001.
- ASIA-AQ Science Team: ASIA-AQ Field Campaign Data, <https://doi.org/doi.org/10.5067/SUBORBITAL/ASIA-AQ/DATA001>, 2024.
- Cambaliza, M. O. L., Shepson, P. B., Bogner, J., Caulton, D. R., Stirr, B., Sweeney, C., Montzka, S. A., Gurney, K. R., Spokas, K., Salmon, O. E., Lavoie, T. N., Hendricks, A., Mays, K., Turnbull, J., Miller, B. R., Lauvaux, T., Davis, K., Karion, A., Moser, B., Miller, C., Obermeyer, C., Whetstone, J., Prasad, K., Miles, N., and Richardson, S.: Quantification and source apportionment of the methane emission flux from the city of Indianapolis, *Elementa: Science of the Anthropocene*, 3, 000037, <https://doi.org/10.12952/journal.elementa.000037>, 2015.
- Cantrell, C. A.: Technical Note: Review of methods for linear least-squares fitting of data and application to atmospheric chemistry problems, *Atmospheric Chemistry and Physics*, 8, 5477–5487, <https://doi.org/10.5194/acp-8-5477-2008>, 2008.
- Crounse, J. D., McKinney, K. A., Kwan, A. J., and Wennberg, P. O.: Measurement of Gas-Phase Hydroperoxides by Chemical Ionization Mass Spectrometry, *Analytical Chemistry*, 78, 6726–6732, <https://doi.org/10.1021/ac0604235>, 2006.
- DiGangi, J. P., Choi, Y., Nowak, J. B., Halliday, H. S., Diskin, G. S., Feng, S., Barkley, Z. R., Lauvaux, T., Pal, S., Davis, K. J., Baier, B. C., and Sweeney, C.: Seasonal Variability in Local Carbon Dioxide Biomass Burning Sources Over Central and Eastern US Using Airborne In Situ Enhancement Ratios, *Journal of Geophysical Research: Atmospheres*, 126, e2020JD034525, <https://doi.org/10.1029/2020JD034525>, 2021.
- DiGangi, J. P., Diskin, G. S., Yoon, S., Alvarez, S. L., Flynn, J. H., Robinson, C. E., Shook, M. A., Thornhill, K. L., Winstead, E. L., Ziemba, L. D., Cambaliza, M. O. L., Simpas, J. B., Hilario, M. R. A., and Sorooshian, A.: Technical note: Apportionment of Southeast Asian Biomass Burning and Urban Influence via In Situ Trace Gas Enhancement Ratios, *EGUsphere*, 2025, 1–16, <https://doi.org/10.5194/egusphere-2025-1454>, 2025.
- Draxler, R. R.: HYSPLIT_4 User's Guide. NOAA technical memorandum ERL ARL-230, 1999.



- 426 Draxler, R. R. and Hess, G. D.: Description of the HYSPLIT4 modeling system, NOAA technical memorandum ERL ARL-
427 224, 1997.
- 428 Draxler, R. R. and Hess, G. D.: An overview of the HYSPLIT_4 modelling system for trajectories, Australian meteorological
429 magazine, 47, 295–308, 1998.
- 430 Fiehn, A., Eckl, M., Kostinek, J., Gałkowski, M., Gerbig, C., Rothe, M., Röckmann, T., Menoud, M., Maazallahi, H., Schmidt,
431 M., Korbeň, P., Nečki, J., Stanisavljević, M., Swolkień, J., Fix, A., and Roiger, A.: Source apportionment of methane emissions
432 from the Upper Silesian Coal Basin using isotopic signatures, *Atmospheric Chemistry and Physics*, 23, 15749–15765,
433 <https://doi.org/10.5194/acp-23-15749-2023>, 2023.
- 434 Guha, A., Gentner, D. R., Weber, R. J., Provencal, R., and Goldstein, A. H.: Source apportionment of methane and nitrous
435 oxide in California’s San Joaquin Valley at CalNex 2010 via positive matrix factorization, *Atmospheric Chemistry and*
436 *Physics*, 15, 12043–12063, <https://doi.org/10.5194/acp-15-12043-2015>, 2015.
- 437 Halliday, H. S., DiGangi, J. P., Choi, Y., Diskin, G. S., Pusede, S. E., Rana, M., Nowak, J. B., Knote, C., Ren, X., He, H.,
438 Dickerson, R. R., and Li, Z.: Using Short-Term CO/CO₂ Ratios to Assess Air Mass Differences Over the Korean Peninsula
439 During KORUS-AQ, *Journal of Geophysical Research: Atmospheres*, 124, 10951–10972,
440 <https://doi.org/10.1029/2018JD029697>, 2019.
- 441 Helfter, C., Tremper, A. H., Halios, C. H., Kotthaus, S., Bjorkegren, A., Grimmond, C. S. B., Barlow, J. F., and Nemitz, E.:
442 Spatial and temporal variability of urban fluxes of methane, carbon monoxide and carbon dioxide above London, UK,
443 *Atmospheric Chemistry and Physics*, 16, 10543–10557, <https://doi.org/10.5194/acp-16-10543-2016>, 2016.
- 444 Hodshire, A. L., Ramnarine, E., Akherati, A., Alvarado, M. L., Farmer, D. K., Jathar, S. H., Kreidenweis, S. M., Lonsdale, C.
445 R., Onasch, T. B., Springston, S. R., Wang, J., Wang, Y., Kleinman, L. I., Sedlacek III, A. J., and Pierce, J. R.: Dilution impacts
446 on smoke aging: evidence in Biomass Burning Observation Project (BBOP) data, *Atmospheric Chemistry and Physics*, 21,
447 6839–6855, <https://doi.org/10.5194/acp-21-6839-2021>, 2021.
- 448 Janhäll, S., Andreae, M. O., and Pöschl, U.: Biomass burning aerosol emissions from vegetation fires: particle number and
449 mass emission factors and size distributions, *Atmospheric Chemistry and Physics*, 10, 1427–1439, <https://doi.org/10.5194/acp-10-1427-2010>, 2010.
- 451 Kurokawa, J. and Ohara, T.: Long-term historical trends in air pollutant emissions in Asia: Regional Emission inventory in
452 ASia (REAS) version 3, *Atmospheric Chemistry and Physics*, 20, 12761–12793, <https://doi.org/10.5194/acp-20-12761-2020>,
453 2020.
- 454 Kurokawa, J., Ohara, T., Morikawa, T., Hanayama, S., Janssens-Maenhout, G., Fukui, T., Kawashima, K., and Akimoto, H.:
455 Emissions of air pollutants and greenhouse gases over Asian regions during 2000–2008: Regional Emission inventory in ASia
456 (REAS) version 2, *Atmospheric Chemistry and Physics*, 13, 11019–11058, <https://doi.org/10.5194/acp-13-11019-2013>, 2013.
- 457 Liang, Y., Weber, R. J., Misztal, P. K., Jen, C. N., and Goldstein, A. H.: Aging of Volatile Organic Compounds in October
458 2017 Northern California Wildfire Plumes, *Environ. Sci. Technol.*, 56, 1557–1567, <https://doi.org/10.1021/acs.est.1c05684>,
459 2022.
- 460 Lin, G. and Wolfe, R.: VIIRS/NPP Imagery Resolution Terrain-Corrected Geolocation L1 6-Min Swath 375m NRT,
461 https://doi.org/10.5067/VIIRS/VNP03IMG_NRT.002, 2022a.
- 462 Lin, G. and Wolfe, R.: VIIRS/NPP Moderate Resolution Terrain-Corrected Geolocation L1 6-Min Swath 750m NRT,
463 https://doi.org/10.5067/VIIRS/VNP03MOD_NRT.002, 2022b.



- 464 Lin, N.-H., Tsay, S.-C., Maring, H. B., Yen, M.-C., Sheu, G.-R., Wang, S.-H., Chi, K. H., Chuang, M.-T., Ou-Yang, C.-F., Fu,
465 J. S., Reid, J. S., Lee, C.-T., Wang, L.-C., Wang, J.-L., Hsu, C. N., Sayer, A. M., Holben, B. N., Chu, Y.-C., Nguyen, X. A.,
466 Sopajaree, K., Chen, S.-J., Cheng, M.-T., Tsuang, B.-J., Tsai, C.-J., Peng, C.-M., Schnell, R. C., Conway, T., Chang, C.-T.,
467 Lin, K.-S., Tsai, Y. I., Lee, W.-J., Chang, S.-C., Liu, J.-J., Chiang, W.-L., Huang, S.-J., Lin, T.-H., and Liu, G.-R.: An overview
468 of regional experiments on biomass burning aerosols and related pollutants in Southeast Asia: From BASE-ASIA and the
469 Dongsha Experiment to 7-SEAS, *Atmospheric Environment*, 78, 1–19, <https://doi.org/10.1016/j.atmosenv.2013.04.066>, 2013.
- 470 Müller, M., Anderson, B. E., Beyersdorf, A. J., Crawford, J. H., Diskin, G. S., Eichler, P., Fried, A., Keutsch, F. N., Mikoviny,
471 T., Thornhill, K. L., Walega, J. G., Weinheimer, A. J., Yang, M., Yokelson, R. J., and Wisthaler, A.: In situ measurements and
472 modeling of reactive trace gases in a small biomass burning plume, *Atmospheric Chemistry and Physics*, 16, 3813–3824,
473 <https://doi.org/10.5194/acp-16-3813-2016>, 2016.
- 474 Nara, H., Tanimoto, H., Tohjima, Y., Mukai, H., Nojiri, Y., and Machida, T.: Emission factors of CO₂, CO and CH₄ from
475 Sumatran peatland fires in 2013 based on shipboard measurements, *Tellus B: Chemical and Physical Meteorology*, 69,
476 1399047, <https://doi.org/10.1080/16000889.2017.1399047>, 2017.
- 477 NASA FIRMS: NRT VIIRS 375 m Active Fire product VNP14IMG, T, 2024.
478 https://doi.org/10.5067/FIRMS/VIIRS/VNP14IMG_NRT.002, 2024.
- 479 NASA VIIRS Characterization Support Team (VCST)/ MODIS Adaptive Processing System (MODAPS): VIIRS/NPP
480 Imagery Resolution 6-Min L1B Swath 375m NRT, https://doi.org/10.5067/VIIRS/VNP02IMG_NRT.002, 2022a.
- 481 NASA VIIRS Characterization Support Team (VCST)/ MODIS Adaptive Processing System (MODAPS): VIIRS/NPP
482 Moderate Resolution Bands L1B 6-Min Swath 750m NRT, https://doi.org/10.5067/VIIRS/VNP02MOD_NRT.002, 2022b.
- 483 Reid, J. S., Hobbs, P. V., Ferek, R. J., Blake, D. R., Martins, J. V., Dunlap, M. R., and Lioussé, C.: Physical, chemical, and
484 optical properties of regional hazes dominated by smoke in Brazil, *Journal of Geophysical Research: Atmospheres*, 103,
485 32059–32080, <https://doi.org/10.1029/98JD00458>, 1998.
- 486 Reid, J. S., Maring, H. B., Narisma, G. T., van den Heever, S., Di Girolamo, L., Ferrare, R., Lawson, P., Mace, G. G., Simpas,
487 J. B., Tanelli, S., Ziemba, L., van Didenhoven, B., Brintjes, R., Bucholtz, A., Cairns, B., Cambaliza, M. O., Chen, G., Diskin,
488 G. S., Flynn, J. H., Hostetler, C. A., Holz, R. E., Lang, T. J., Schmidt, K. S., Smith, G., Sorooshian, A., Thompson, E. J.,
489 Thornhill, K. L., Trepte, C., Wang, J., Woods, S., Yoon, S., Alexandrov, M., Alvarez, S., Amiot, C. G., Bennett, J. R., Brooks,
490 M., Burton, S. P., Cayan, E., Chen, H., Collow, A., Crosbie, E., DaSilva, A., DiGangi, J. P., Flagg, D. D., Freeman, S. W.,
491 Fu, D., Fukada, E., Hilario, M. R. A., Hong, Y., Hristova-Veleva, S. M., Kuehn, R., Kowch, R. S., Leung, G. R., Loveridge,
492 J., Meyer, K., Miller, R. M., Montes, M. J., Moum, J. N., Nenes, A., Nesbitt, S. W., Norgren, M., Nowotnick, E. P., Rauber,
493 R. M., Reid, E. A., Rutledge, S., Schlosser, J. S., Sekiyama, T. T., Shook, M. A., Sokolowsky, G. A., Stammes, S. A., Tanaka,
494 T. Y., Wasilewski, A., Xian, P., Xiao, Q., Xu, Z., and Zavaleta, J.: The Coupling Between Tropical Meteorology, Aerosol
495 Lifecycle, Convection, and Radiation during the Cloud, Aerosol and Monsoon Processes Philippines Experiment (CAMP2Ex),
496 *Bulletin of the American Meteorological Society*, 104, E1179–E1205, <https://doi.org/10.1175/BAMS-D-21-0285.1>, 2023.
- 497 Reinecke, T., Leiminger, M., Jordan, A., Wisthaler, A., and Müller, M.: Ultrahigh Sensitivity PTR-MS Instrument with a Well-
498 Defined Ion Chemistry, *Analytical Chemistry*, 95, 11879–11884, <https://doi.org/10.1021/acs.analchem.3c02669>, 2023.
- 499 Sachse, G. W., Hill, G. F., Wade, L. O., and Perry, M. G.: Fast-response, high-precision carbon monoxide sensor using a
500 tunable diode laser absorption technique, *Journal of Geophysical Research: Atmospheres*, 92, 2071–2081,
501 <https://doi.org/10.1029/JD092iD02p02071>, 1987.



- 502 Sachse, G. W., Jr., J. E. C., Hill, G. F., Wade, L. O., Burney, L. G., and Ritter, J. A.: Airborne tunable diode laser sensor for
503 high-precision concentration and flux measurements of carbon monoxide and methane, in: *Proc.SPIE*, 157–166,
504 <https://doi.org/10.1117/12.46162>, 1991.
- 505 Schroeder, W., Oliva, P., Giglio, L., and Csizsar, I. A.: The New VIIRS 375m active fire detection data product: Algorithm
506 description and initial assessment, *Remote Sensing of Environment*, 143, 85–96, <https://doi.org/10.1016/j.rse.2013.12.008>,
507 2014.
- 508 Stein, A. F., Draxler, R. R., Rolph, G. D., Stunder, B. J. B., Cohen, M. D., and Ngan, F.: NOAA’s HYSPLIT Atmospheric
509 Transport and Dispersion Modeling System, *Bulletin of the American Meteorological Society*, 96, 2059–2077,
510 <https://doi.org/10.1175/BAMS-D-14-00110.1>, 2015.
- 511 Torres, O. and Goddard Earth Sciences Data and Information Services Center (GES DISC): OMPS-NPP L2 NM Aerosol Index
512 swath orbital V2, <https://doi.org/10.5067/40L92G8144IV>, 2019.
- 513 Tsay, S.-C., Hsu, N. C., Lau, W. K.-M., Li, C., Gabriel, P. M., Ji, Q., Holben, B. N., Judd Welton, E., Nguyen, A. X., Janjai,
514 S., Lin, N.-H., Reid, J. S., Boonjawat, J., Howell, S. G., Huebert, B. J., Fu, J. S., Hansell, R. A., Sayer, A. M., Gautam, R.,
515 Wang, S.-H., Goodloe, C. S., Miko, L. R., Shu, P. K., Loftus, A. M., Huang, J., Kim, J. Y., Jeong, M.-J., and Pantina, P.: From
516 BASE-ASIA toward 7-SEAS: A satellite-surface perspective of boreal spring biomass-burning aerosols and clouds in
517 Southeast Asia, *Atmospheric Environment*, 78, 20–34, <https://doi.org/10.1016/j.atmosenv.2012.12.013>, 2013.
- 518 Vay, S. A., Woo, J.-H., Anderson, B. E., Thornhill, K. L., Blake, D. R., Westberg, D. J., Kiley, C. M., Avery, M. A., Sachse,
519 G. W., Streets, D. G., Tsutsumi, Y., and Nolf, S. R.: Influence of regional-scale anthropogenic emissions on CO₂ distributions
520 over the western North Pacific, *Journal of Geophysical Research: Atmospheres*, 108, <https://doi.org/10.1029/2002JD003094>,
521 2003.
- 522 Warneke, C., Schwarz, J. P., Dibb, J., Kalashnikova, O., Frost, G., Al-Saad, J., Brown, S. S., Brewer, Wm. A., Soja, A., Seidel,
523 F. C., Washenfelder, R. A., Wiggins, E. B., Moore, R. H., Anderson, B. E., Jordan, C., Yacovitch, T. I., Herndon, S. C., Liu,
524 S., Kuwayama, T., Jaffe, D., Johnston, N., Selimovic, V., Yokelson, R., Giles, D. M., Holben, B. N., Goloub, P., Popovici, I.,
525 Trainer, M., Kumar, A., Pierce, R. B., Fahey, D., Roberts, J., Gargulinski, E. M., Peterson, D. A., Ye, X., Thapa, L. H., Saide,
526 P. E., Fite, C. H., Holmes, C. D., Wang, S., Coggon, M. M., Decker, Z. C. J., Stockwell, C. E., Xu, L., Gkatzelis, G., Aikin,
527 K., Lefer, B., Kaspari, J., Griffin, D., Zeng, L., Weber, R., Hastings, M., Chai, J., Wolfe, G. M., Hanisco, T. F., Liao, J.,
528 Campuzano Jost, P., Guo, H., Jimenez, J. L., Crawford, J., and The FIREX-AQ Science Team: Fire Influence on Regional to
529 Global Environments and Air Quality (FIREX-AQ), *Journal of Geophysical Research: Atmospheres*, 128, e2022JD037758,
530 <https://doi.org/10.1029/2022JD037758>, 2023.
- 531 Yokelson, R. J., Griffith, D. W. T., and Ward, D. E.: Open-path Fourier transform infrared studies of large-scale laboratory
532 biomass fires, *Journal of Geophysical Research: Atmospheres*, 101, 21067–21080, <https://doi.org/10.1029/96JD01800>, 1996.
- 533 Yokelson, R. J., Crounse, J. D., DeCarlo, P. F., Karl, T., Urbanski, S., Atlas, E., Campos, T., Shinozuka, Y., Kapustin, V.,
534 Clarke, A. D., Weinheimer, A., Knapp, D. J., Montzka, D. D., Holloway, J., Weibring, P., Flocke, F., Zheng, W., Toohey, D.,
535 Wennberg, P. O., Wiedinmyer, C., Mauldin, L., Fried, A., Richter, D., Walega, J., Jimenez, J. L., Adachi, K., Buseck, P. R.,
536 Hall, S. R., and Shetter, R.: Emissions from biomass burning in the Yucatan, *Atmospheric Chemistry and Physics*, 9, 5785–
537 5812, <https://doi.org/10.5194/acp-9-5785-2009>, 2009.
- 538 Yokelson, R. J., Andreae, M. O., and Akagi, S. K.: Pitfalls with the use of enhancement ratios or normalized excess mixing
539 ratios measured in plumes to characterize pollution sources and aging, *Atmospheric Measurement Techniques*, 6, 2155–2158,
540 <https://doi.org/10.5194/amt-6-2155-2013>, 2013.



541 York, D., Evensen, N. M., Martínez, M. L., and De Basabe Delgado, J.: Unified equations for the slope, intercept, and standard
542 errors of the best straight line, American Journal of Physics, 72, 367–375, <https://doi.org/10.1119/1.1632486>, 2004.

543