

Technical Note: Identifying biomass burning emissions during ASIA-AQ using greenhouse gas enhancement ratios

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Abstract. Biomass burning (BB) is a primary source of atmospheric chemistry reactants, aerosols, and greenhouse gases. Smoke plumes have air quality impacts local to the fire itself and regionally via long distance transport. Open burning of agriculture fields in Southeast Asia leads to frequent seasonal occurrences of regional BB-induced smoke haze and long-range transport of BB particles via the northeast monsoon. The Airborne and Satellite Investigation of Asian Air Quality (ASIA-AQ) campaign visited several areas including the Philippines, South Korea, Thailand, and Taiwan during a time of agricultural burning. This campaign consisted of airborne measurements on the NASA DC-8 aircraft aimed to validate observations from South Korea's Geostationary Environment Monitoring Spectrometer (GEMS) and to address local air quality challenges. We developed a method that used a combination of BB markers to identify ASIA-AQ DC-8 data influenced by BB and flag them for further analysis. Specifically, we used rolling slope enhancement ratios of CO/CO₂ and CH₄/CO along with mixing ratios of CH₃CN, HCN, and CO, and particle scattering coefficient measurements. The flag was triggered when a combination of these variables exceeded a flight specific threshold. We found varying levels of BB-influence in the areas studied, with data flagged for BB being <1% for the Philippines and Korea, and <2% for Taiwan, but 19% for Thailand. Our method for flagging ASIA-AQ BB-affected data can be used to focus additional analyses of the ASIA-AQ campaign such as pairing with back trajectories, satellite hotspot products, and microphysical aerosol characteristics.

1 Introduction

As Asian economies and populations continue to grow, so will their contribution to global greenhouse gas (GHG) emissions, driven primarily by increases in fossil fuel combustion and seasonally by biomass burning (BB) events. Left unchecked, these emissions will negatively impact air quality and climate, and therefore it is imperative that emission sources are properly identified and accounted for in emission inventories. Sources of methane (CH₄) can include: fossil fuels (coal mining),

34 agricultural emissions (enteric fermentation and rice cultivation), and solid waste disposal and wastewater treatment; and for
 35 carbon dioxide (CO₂): fossil fuel/biofuel combustion processes (powerplants and transportation) and industrial processes
 36 (cement production) (Kurokawa et al., 2013; Kurokawa and Ohara, 2020). Carbon monoxide (CO), while not a greenhouse
 37 gas is also emitted from urban sources such as domestic (residential heating and cooking), industrial, and transport sectors
 38 (Kurokawa et al., 2013; Kurokawa and Ohara, 2020). Biomass burning is also a significant source of these GHGs and CO;
 39 however, the seasonal nature of crop residue burning, and unpredictability of natural fires makes it difficult to accurately
 40 quantify these variables sources (Akagi et al., 2011; Streets et al., 2003; Crutzen and Andreae, 1990).

41 The Airborne and Satellite Investigation of Asian Air Quality (ASIA-AQ) field campaign was conducted during February-
 42 March 2024 near the maximum period of seasonal agricultural burning in Southeast Asia, providing an ideal opportunity to
 43 further study biomass burning. ASIA-AQ was an international joint air quality campaign between the National Aeronautics
 44 and Space Administration (NASA) and several space and environmental agencies in Korea (National Institute of
 45 Environmental Research (NIER) and Korea Meteorological Administration (KMA)), the Philippines (Department of
 46 Environment and Natural Resources (DENR), Philippine Space Agency (PhilSA), and Manila Observatory), Thailand (Geo-
 47 Informatics and Space Technology Development Agency (GISTDA) and Pollution Control Department (PCD)), and Taiwan
 48 (Ministry of Environment, National Central University (NCU), and Academia Sinica). Observations included in-situ
 49 measurements of trace gasses, aerosol properties, radiation fluxes, and meteorological parameters from NASA's DC-8 aircraft
 50 and several Korean aircraft, remote sensing measurements from NASA's G-III aircraft, satellite observations from Korea's
 51 GEMS instrument, and several ground station measurements.

52 Current observation-based (top-down) techniques for GHG source apportionment include aircraft-based mass balances
 53 (Cambaliza et al., 2015), standard eddy covariance (Helfter et al., 2016), positive matrix factorization (Guha et al., 2015), and
 54 isotopic signatures (Fiehn et al., 2023). Enhancement ratios between GHGs have also been used to characterize sources of BB
 55 plumes (Akagi et al., 2011; Andreae and Merlet, 2001; Nara et al., 2017; Yokelson et al., 1996); however, Yokelson et al.
 56 (2013) pointed out several issues with this technique including problems that arise when background concentrations are
 57 changing over the measurement period. Halliday et al. (2019) used short-term $\Delta\text{CO}/\Delta\text{CO}_2$ calculated over 60s rolling windows
 58 and filtered data by the coefficient of determination (R^2) to negate the need for a consistent background measurement. DiGangi
 59 et al. (2021) applied a similar method to data from NASA's Atmospheric Carbon Transport-America (ACT-America)
 60 campaign to characterize local CO₂ emissions. DiGangi et al. (2025) adapted this method for use with CH₄/CO ratios for
 61 NASA's Cloud, Aerosol, and Monsoon Processes Philippines Experiment (CAMP²EX) airborne field campaign in a chemical
 62 influence flag, with a focus on separating BB and urban influences. Previous work during the Fire Influence on Regional to
 63 Global Environments and Air Quality (FIREX-AQ) airborne campaign used CO and black carbon (BC) enhancements to
 64 identify and flag smoke plumes (Warneke et al., 2023). Since the ASIA-AQ campaign overflowed many urban areas, a more BB-
 65 specific approach was required. The method presented in this work combines the use of CO/CO₂ and CH₄/CO ratios to further
 66 constrain source contributions. To specifically target BB sources, we have also incorporated mixing ratios of acetonitrile
 67 (CH₃CN) (Yokelson et al., 2009; Lobert et al., 1990; Holzinger et al., 1999), hydrogen cyanide (HCN) (Yokelson et al., 2009;

68 Lobert et al., 1990; Holzinger et al., 1999), CO (Lin et al., 2013; Nara et al., 2017), and aerosol scattering coefficient (Lin et
69 al., 2013; Tsay et al., 2013) into our method. This work provides a targeted method to identify air masses influenced by biomass
70 burning during the 2024 ASIA-AQ airborne field campaign, an example of which is shown in our Thailand case study.

71 **2 Data and Methods**

72 **2.1 ASIA-AQ Field Campaign and In Situ Aircraft Measurements**

73 These methods were developed for the ASIA-AQ campaign, which was a joint international field study focused on air quality
74 challenges local to its areas of study, which included the Philippines, Korea, Thailand, and Taiwan. Data for this campaign
75 were collected from aircraft, including NASA's DC-8 and G-III and three Korean aircraft, and several ground sites. Sixteen
76 DC-8 science flights were conducted between 6 February 2024 and 27 March 2024; the full flight break-down can be found
77 in Table S1 along with typical flight paths in Figures S1-4. The NASA DC-8 aircraft carried a variety of instruments used for
78 in-situ gas-phase, aerosol, radiation, and meteorological measurements. CO₂ was measured via non-dispersive infrared
79 spectroscopy using a modified LICOR 7000 instrument with an uncertainty of 0.25 ppm at <500 ppm and 2% at > 500 ppm
80 (Vay et al., 2003). CH₄ and CO were measured via wavelength modulation spectroscopy using the Differential Absorption
81 Carbon monoxide Measurement (DACOM) instrument, with CO uncertainties of 2% for < 1 ppm and 5% for > 1 ppm and
82 CH₄ uncertainty of 1% (Sachse et al., 1987, 1991). HCN mixing ratios were measured using a Chemical Ionization High
83 Resolution Time of Flight Mass Spectrometer with an uncertainty of 25% plus 35 pptv (CIT-CIMS) (Crounse et al., 2006).
84 CH₃CN measurements were taken using a Proton-Transfer-Reaction Time-of-Flight Mass Spectrometer (PTR-MS) with an
85 uncertainty of 13% (Müller et al., 2016; Reinecke et al., 2023). Total aerosol scattering at 550 nm measurements were measured
86 using a TSI-3563 Nephelometer with an estimated accuracy of 20% and estimated precision of 1 Mm⁻¹ (Anderson and Ogren,
87 1998). Accumulation-mode aerosol BC mass concentrations were measured by a Single Particle Soot Photometer (SP2-D)
88 (Droplet Measurement Technologies Inc. (DMT), Longmont, Colorado, USA) with an uncertainty of 20%, operated by NASA
89 Langley Research Center. Optical particle size distributions for particles with diameters between 63.1 and 1000 nm were
90 measured using a DMT Ultra High Sensitivity Aerosol Spectrometer (UHSAS) Model UHSAS-0.055 (DMT, Longmont
91 Colorado USA) with an uncertainty of 20%, while particles with diameters between 3.16 and 89.1 nm were sized using a
92 Scanning Mobility Particle Sizer (SMPS) with an uncertainty of 20%. Submicron particle number concentrations (>10 nm)
93 were measured with a TSI Condensation Particle Counter (CPC) CPC-3772 (TSI Inc., Shoreview, Minnesota, USA) with an
94 uncertainty of 10%. A duplicate instrument also with 10% uncertainty was used to measure non-volatile particle number
95 concentrations (>10 nm) with the sample passing through a thermal denuder heated to 350 °C prior to entering the instrument.
96 All DC-8 data used in this work had a time resolution of 1 Hz, with the exceptions of the SMPS which measured every minute.
97 Hourly Thai precipitation measurements were obtained from the Thailand Pollution Control Department using the Phaya Thai,
98 Khet Phaya Thai, and Bangkok sites. This data and the developed BB flag are publicly available on the ASIA-AQ data archive
99 (ASIA-AQ Science Team, 2024).

100 2.2 Rolling Slope Method

101 The rolling slope calculation method employed here was taken from Halliday et al. (2019), where linear fits over 120s rolling
102 windows of CO vs. CO₂ and CH₄ vs. CO were calculated using error adjusted bivariate regression as detailed in York et al.
103 (2004) and Cantrell (2008). Applying this method, a maximum of 121 observations were used in each calculation, while the
104 minimum for a valid calculation was set at three. Slopes with low goodness of fit ($R^2 < 0.5$) were filtered out as uncorrelated,
105 and slopes with a ΔCH_4 , ΔCO , or ΔCO_2 less than five times the precision value were dropped. While Halliday et al. (2019)
106 demonstrated that the shapes of the slope distributions are somewhat insensitive to the R^2 cutoff (when $R^2 \geq 0.5$), here we
107 observed differences between correlated slopes ($R^2 \geq 0.5$) and uncorrelated slopes ($R^2 < 0.5$). Halliday et al. (2019) also showed
108 that varying the window width did not drastically change the slope distributions. Similarly, Figures S5-7 show that with
109 increasing window width, the number of correlated slopes increases and with an increasing R^2 cutoff the number decreases;
110 however, the shapes of the slope distributions are similar. Table S2 exhibits the percentage of calculated $\Delta\text{CO}/\Delta\text{CO}_2$ and
111 $\Delta\text{CH}_4/\Delta\text{CO}$ slopes that were correlated for each flight. On average across all flights, 52% of $\Delta\text{CO}/\Delta\text{CO}_2$ slopes and 53% of
112 $\Delta\text{CH}_4/\Delta\text{CO}$ slopes were correlated according to the $R^2 \geq 0.5$ criterion, in 120s rolling windows.

113

114 $\Delta\text{CO}/\Delta\text{CO}_2$ rolling slopes can be used to isolate plumes or air mass boundaries with correlated behavior and provide a
115 rudimentary source classification. The combustion efficiency of the source is inversely related to the $\Delta\text{CO}/\Delta\text{CO}_2$ slope,
116 therefore we have divided up source behaviors into four combustion efficiency bins based on $\Delta\text{CO}/\Delta\text{CO}_2$ slopes: 0-1%, 1-2%,
117 2-4%, and >4%, as previously used by Halliday et al., (2019). In this work, the >4% combustion efficiency bin is most relevant
118 for identifying BB-influenced plumes. Figure 1A shows an example of this classification applied to the CO and CO₂ bulk
119 ratios from Thailand Flight 1 demonstrating the ability of this method to highlight different air mass behaviors. $\Delta\text{CH}_4/\Delta\text{CO}$
120 slopes have been used to distinguish between urban sources of methane and biomass burning (DiGangi et al., 2025; Reid et
121 al., 2023), with higher slopes indicative of more urban sources and smaller slopes $0 < x < \sim 40\%$ with biomass burning. Figure
122 1B displays the application of these regimes to CO and CH₄ bulk ratios. Both Fig. 1A and 1B demonstrate a large influence
123 from slopes indicative of biomass burning, >4% for $\Delta\text{CO}/\Delta\text{CO}_2$ (shown in dark red) and $0 < x < \sim 40\%$ for $\Delta\text{CH}_4/\Delta\text{CO}$ (shown
124 in orange).

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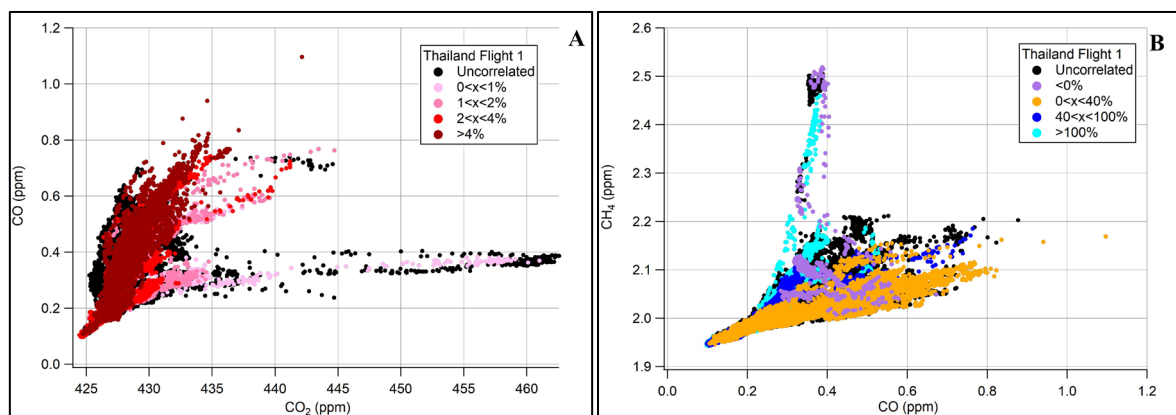


Figure 1. CO vs CO₂ mixing ratios colored by $\Delta\text{CO}/\Delta\text{CO}_2$ rolling slope enhancement ratio regimes (A) and CH₄ vs CO mixing ratios colored by $\Delta\text{CH}_4/\Delta\text{CO}$ rolling slope enhancement ratio regimes (B) for Thailand flight 1.

2.3 Biomass Burning Flag Determination

The biomass burning flag uses a combination of variables indicative of or a product of biomass burning, including CO mixing ratio, particle scattering at 550 nm, HCN, CH₃CN, $\Delta\text{CO}/\Delta\text{CO}_2$, and $\Delta\text{CH}_4/\Delta\text{CO}$. The BB flag is triggered when at least two of the variables exceed their flight-specific thresholds, except for $\Delta\text{CH}_4/\Delta\text{CO}$ which must fall in a range between zero and its threshold (Table S3). Using a single variable was deemed insufficient due to the possibility of confounding factors, for example, the utility of CH₃CN as a BB maker has been shown to be less effective in urban areas due to interference from vehicle and solvent usage emissions (Huangfu et al., 2021). However, when CO mixing ratio and particle scattering are paired together a third variable needs to meet its threshold to trigger the flag; the reasoning for this is to prevent false positives occurring from combustion processes other than biomass burning, such as transportation and industrial sources. Additionally, the BB flag is triggered when all four non-slope variables are within 10% of their threshold to address edge cases. Example scenarios for triggering the flag are shown in Table 1.

150 Table 1. Example scenarios for triggering the BB flag for Thailand flight 1 (03/16/2024). Underlined values represent the variables
 151 that have exceeded their threshold (x), italic values are variables within 10% of their threshold (x). The last row represents an
 152 example where all four non-slope variables are at least within 10% of their threshold (x).

BB Flag	$0 < \Delta\text{CH}_4/\Delta\text{CO} < x$ (%)	$\Delta\text{CO}/\Delta\text{CO}_2 > x$ (%)	$\text{CO} > x$ (ppm)	$\text{HCN} > x$ (ppt)	$\text{CH}_3\text{CN} > x$ (ppb)	Total scattering @ 550 nm $> x$ (Mm^{-1})
Threshold (x)	44.5	4	0.32	2750	1.50	160
✓	N/A	<u>5.1</u>	<u>0.67</u>	<u>2810</u>	1.31	<u>487</u>
✓	<u>40.3</u>	<u>9.8</u>	<u>0.34</u>	1310	0.56	<u>227</u>
✗	54.6	2.5	0.22	774	0.30	36
✓	N/A	N/A	<u>0.53</u>	<u>2823</u>	0.91	<u>386</u>
✗	N/A	N/A	<u>0.33</u>	1053	0.57	<u>186</u>
✓	N/A	2.2	<u>0.67</u>	<i>2715</i>	<i>1.43</i>	<u>418</u>

153 2.3.1 Variable Threshold Determination

154 The BB flag shares similarities with the Chemical Influence flag developed for CAMP²EX, where lower positive $\Delta\text{CH}_4/\Delta\text{CO}$
 155 rolling slopes were associated with biomass burning influence (DiGangi et al., 2025; Reid et al., 2023). Here, the specific
 156 $\Delta\text{CH}_4/\Delta\text{CO}$ cutoff values were chosen on a flight-by-flight basis by evaluating the slope distribution for each flight and looking
 157 for distinctive populations. An example of this process is shown in Fig. 2A, where there is a distribution separation occurring
 158 at a slope value of 44.5% for the first Thailand flight. For $\Delta\text{CO}/\Delta\text{CO}_2$ the threshold was set to greater than 4% for all flights,
 159 as low efficiency processes like biomass burning are associated with higher $\Delta\text{CO}/\Delta\text{CO}_2$ values (Halliday et al., 2019). The
 160 thresholds for the other variables were initially determined by identifying at which concentration within the biomass burning
 161 regime (as set by $\Delta\text{CH}_4/\Delta\text{CO}$) the variable is distinguishable from non-biomass burning influence regimes. In the Fig. 2B
 162 example, that regime is from 0-44.5% $\Delta\text{CH}_4/\Delta\text{CO}$, as determined from Fig. 2A, and the CH_3CN threshold was set to 1.00 ppb.
 163 This method was repeated for the other variables (HCN, particle scattering, and CO mixing ratio) and for all flights. Thresholds
 164 for the mixing ratio and scattering variables were set on a flight-by-flight basis. Flight specific thresholds were chosen over
 165 campaign thresholds due to changing background concentrations, time periods, and environmental conditions for each flight
 166 and location.
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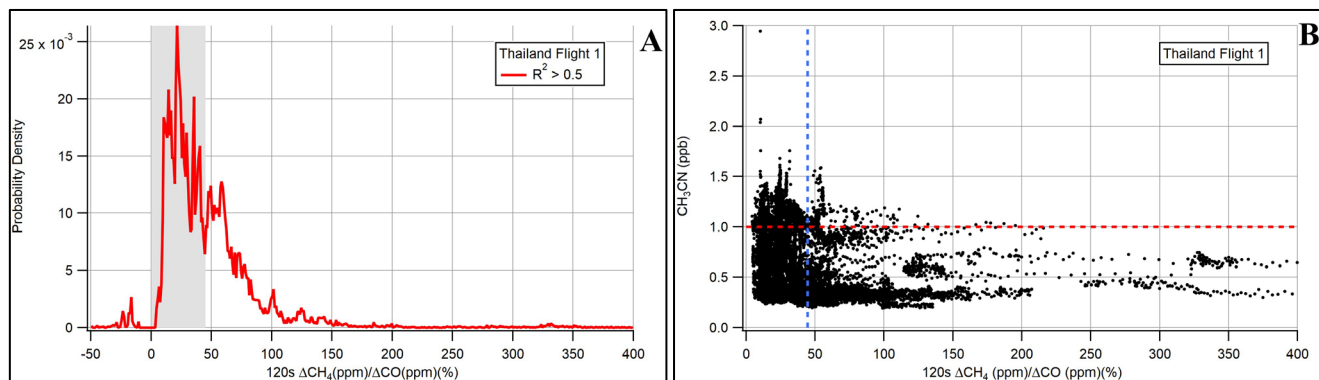


Figure 2 (A) Frequency distribution of rolling slope $\Delta\text{CH}_4/\Delta\text{CO}$ enhancement ratios. The shaded region is the set BB regime for this flight, as determined by the minimum in the distribution. (B) CH_3CN vs $\Delta\text{CH}_4/\Delta\text{CO}$ for Thailand flight 1. The blue vertical line represents the edge of the biomass burning influence regime at 44.5% $\Delta\text{CH}_4/\Delta\text{CO}$, while the red horizontal line is the preliminary CH_3CN threshold for the BB flag. The upper left quadrant of the figure represents the BB regime.

In the case of thresholds that were more difficult to distinguish, a higher value was initially selected to minimize false positives in the flag. Thresholds were then further refined from this initial value by examining the sensitivity of the total percentage of points flagged to that threshold. The final thresholds were chosen at the lowest levels for which the percentage of points remained essentially constant. Figure 3 demonstrates the results from one of these sensitivity tests performed on the first flight in Thailand. In this case, decreasing the thresholds for CO mixing ratio, CH_3CN , HCN, and particle scattering results in large increases in the number of points flagged. In Fig. 3A, the dashed HCN and CH_3CN traces are still decreasing past their initial threshold values represented by the triangles (2400 ppt and 1.0 ppb); therefore, this is motivation to increase their thresholds by ~15% and ~50% respectively, resulting in the solid traces. The initial thresholds for CO mixing ratio and particle scattering (0.32 ppm and 160), shown as triangles Fig. 3B required no adjustment for this particular flight; the decrease in the percentage of points flagged for the post-adjustment traces is due to the increases in the HCN and CH_3CN thresholds. The final determined thresholds and rolling slope regimes for all flights are in Table S3, while Table 1 provides some example scenarios of the BB flag being triggered or not for Thailand flight 1. The concentration thresholds varied by a factor of 2.6 for CO and up to 8 for CH_3CN , this variability can be explained by changing background concentrations for each flight day and location. This approach was followed to minimize false positives in the dataset; therefore, this method is less sensitive to air masses with only minor influences from biomass burning.

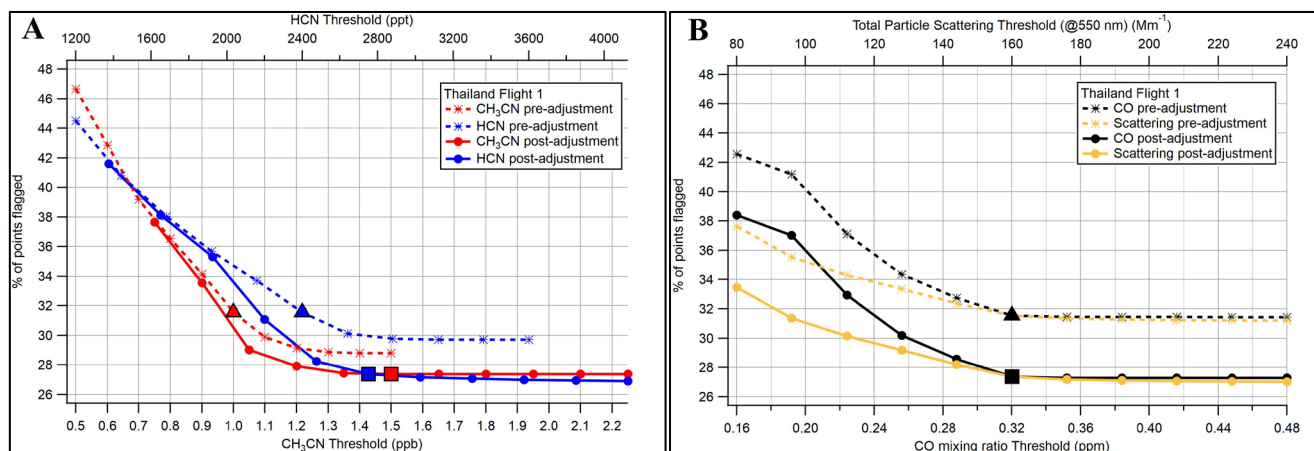


Figure 3. Example variable threshold sensitivity plots for the first Thailand flight. Fig 3A shows the sensitivities for the HCN and CH₃CN thresholds established using plots like Fig 2B in the dashed traces and the final thresholds in solid. Fig 3B shows the sensitivities for the CO mixing ratio and total particle scattering thresholds established using plots like Fig 2B in the dashed traces and the final thresholds in solid. The triangles represent the initial thresholds for each variable and the squares the final thresholds.

2.4 HYSPLIT Back Trajectories

Air mass history was probed using 48 hour back trajectories calculated from the DC-8 flight track at one second intervals using NOAA's HYSPLIT model using GFS 0.25° meteorology (Draxler, 1999; Draxler and Hess, 1997, 1998; Stein et al., 2015). For more specific source type, altitude, and receptor location analysis, we used the BB flag as a filter and focused on specific areas and altitudes of the flight path, such as urban areas at low altitudes to determine whether the emissions measured in these locations were local or transport. To empirically quantify these contributions, we totaled the number of back trajectory points under 1 km that traveled through a certain area prior to ending up at the receptor location; locations used in this analysis are described in Table S4 and shown in Figure S8.

2.5 VIIRS Fire Hotspots & Imagery

The Visible Infrared Imaging Radiometer Suite (VIIRS) I-Band 375 m Active Fire data product was used to assess fire hotspot density over Southeast Asia (Schroeder et al., 2014). Data version 2.0 from the Suomi National Polar-Orbiting Partnership (Suomi NPP) spacecraft were downloaded from NASA-FIRMS (Fire Information for Resource Management System) (NASA FIRMS, 2024). VIIRS Corrected Reflectance (True Color) imagery was downloaded from NASA-FIRMS (Lin and Wolfe, 2022a, b; NASA VIIRS Characterization Support Team (VCST)/ MODIS Adaptive Processing System (MODAPS), 2022a, b). The Ozone Mapping and Profiling Suite (OMPS) Aerosol Index overlay was downloaded from NASA-FIRMS (Torres and Goddard Earth Sciences Data and Information Services Center (GES DISC), 2019).

209 **3 Results and Discussion**

210 **3.1 Breakdown of points flagged for each area**

211 Figure 4A breaks down the percentage of points flagged for BB for each flight and for each location, with a breakdown based
212 on cities in Fig 4B. Very little BB-influence was observed in the Philippines, with most of it occurring north of Manila under
213 1.5 km above ground level (AGL) (Fig S1). For Korea, the highest occurrences of BB occurred on flights 2 and 3. On flight 2
214 the majority of the BB was observed over the West Sea under 500 m AGL. For flight 3 there was BB-influence over both the
215 West Sea and Seoul above 1 km AGL (Fig S2). Most of the observed BB-influence for the campaign occurred in Thailand,
216 specifically flights 1 and 2, with Chiang Mai experiencing more BB compared to Bangkok (Fig S3), except for flight 3. Section
217 3.2 delves more into the geographical and temporal BB flag differences observed in Thailand during the campaign. Taiwan
218 had the next highest percent occurrence, but Korea overall had more points flagged for BB compared to Taiwan. The total time
219 spent collecting data over these locations is the reason for this discrepancy, the campaign spent almost four times as much time
220 over Korea compared to Taiwan. However, in terms of cities, Kaohsiung had the third highest percentage and number of points
221 flagged, after Chiang Mai and Bangkok, with most of these flagged points occurring on flight 1. Most of the BB observed in
222 Taiwan occurred during the third flight, with the majority occurring on parts of the flight track not over Kaohsiung and above
223 1 km AGL (Fig S4).

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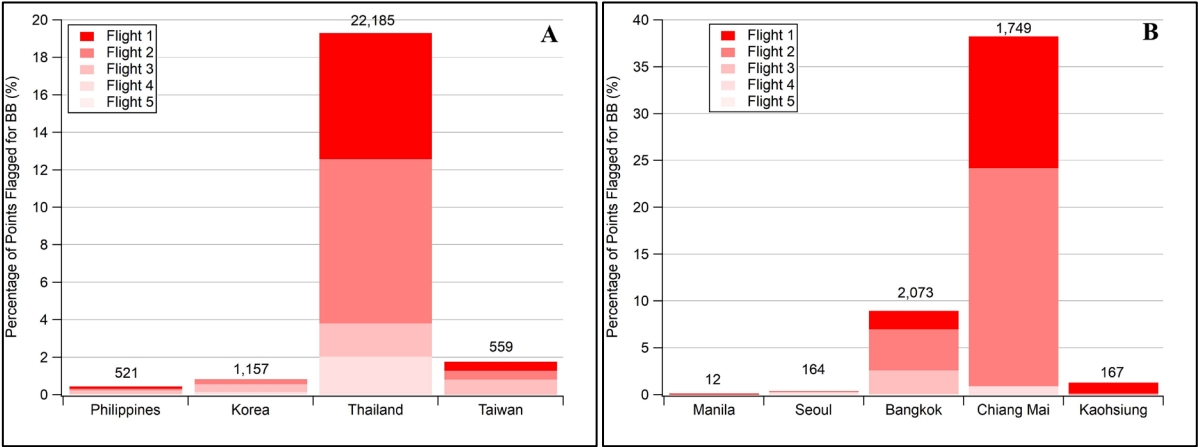
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226 **Figure 4. Percentage of points flagged for BB for each flight over each country (A) and select cities (B) with total number of points**
227 **flagged displayed.**

228 **3.2 Thailand Case-Study**

229 **3.2.1 Flight by Flight Breakdown**

230 Since Thailand had the largest prevalence of biomass burning among the flight locations, we have broken down these flights
231 in greater detail. Figure 5 shows the flight-by-flight breakdown of the BB flag (A-D), VIIRS 48h fire hotspot density maps (E-



H) and VIIRS satellite imagery with the OMPS aerosol index overlay (I-L) for the four Thailand flights. While flight 2 had the highest overall occurrence of BB (10,092 points compared to 7,725) (Fig 5B), flight 1 had more BB-influence closer to the surface (1,289 compared to 1,144) (Fig 5A). An increase in fire hotspot density is also observed when comparing Fig 5E to 5F, where the density has increased in Eastern Thailand, Cambodia, and Northern Thailand/Myanmar. However, a similar increase is not observed looking at the OMPS aerosol index overlays (Fig 5I to Fig 5J). For flight 1 (Fig 5I) the aerosol index was elevated over northern Laos, northern Vietnam, northern and western Thailand, and central and eastern Myanmar, indicating either a more concentrated, thicker and/or higher layer of absorbing aerosols (dust and smoke) in these areas. For flight 2 (Fig 5J) the higher aerosol index values were concentrated over Hanoi, with western Thailand clearing up and few elevated values over eastern Thailand, Cambodia, and southern Laos. This apparent aerosol index decrease around Thailand contrasts with what was observed using the BB flag and the VIIRS fire hotspot density maps, however, the OMPS aerosol index is also sensitive to changes in aerosol height, concentration, and degree of absorption (Torres and Goddard Earth Sciences Data and Information Services Center (GES DISC), 2019). Prior to flight 3, the area received a considerable amount of rain (~41 mm in Bangkok and up to 82 mm on the northern flight track on the day before flight 3), which likely washed out most of the smoke and put out local fires, seen in the decreased fire hotspot density across Thailand (Fig 5G), and decreased aerosol indices (Fig 5K). For the final flight, there was a slight increase in the number of points flagged compared to flight 3 (~17%), more near the surface as burning resumed, which is apparent when comparing the VIIRS fire hotspot density maps (Fig 5G and Fig 5H) and scattered elevated aerosol indices over northern Thailand (Fig 5L).

Figure 5A-D also shows a breakdown of BB-influence in terms of major population centers on the flight track, specifically Bangkok on the southern portion of the track and Chiang Mai on the northern part. For flight 1, no BB-influence was observed below 1 km in Bangkok across four low passes. This contrasts with observations during one pass in Chiang Mai where ~80% of the low altitude data were flagged for BB. This is consistent with Fig 5E, where the fire hot spot densities were higher around Chiang Mai compared to Bangkok, and Fig 5I with elevated aerosol indices over Chiang Mai. For flight 2, there were equivalent amounts of BB-influence over the two cities, but significantly more at the surface in Chiang Mai, however, relative to the time spent at each location over 90% of high-level data from Chiang Mai were flagged while only ~25% of high-level Bangkok data were. Even though the fire hotspot density around Bangkok was higher for flight 2 (Fig 5F) compared to flight 1 (Fig 5E), that BB-influence does not appear to have reached the surface. One possible explanation is that the prevailing southerly winds were blowing that smoke more towards Central Thailand. In the third flight, the rain likely washed out almost all of the smoke around Chiang Mai and quelled many fires as observed in Fig 5G. However, there was still some BB-influence observed above Bangkok, perhaps by a transported airmass aloft unaffected by the rain. From Fig 5G, there were still high fire hotspot densities occurring in Cambodia and Laos, which could be traced forward to Bangkok using HYSPLIT back trajectories (Fig 6A). For the last flight, there was little BB-influence observed in both cities (~1% total) but still a meaningful amount across the rest of the flight track (~8%). Figure 5H shows that fire activities seemed to have resumed throughout Thailand in the 48 hours leading up to flight 4.

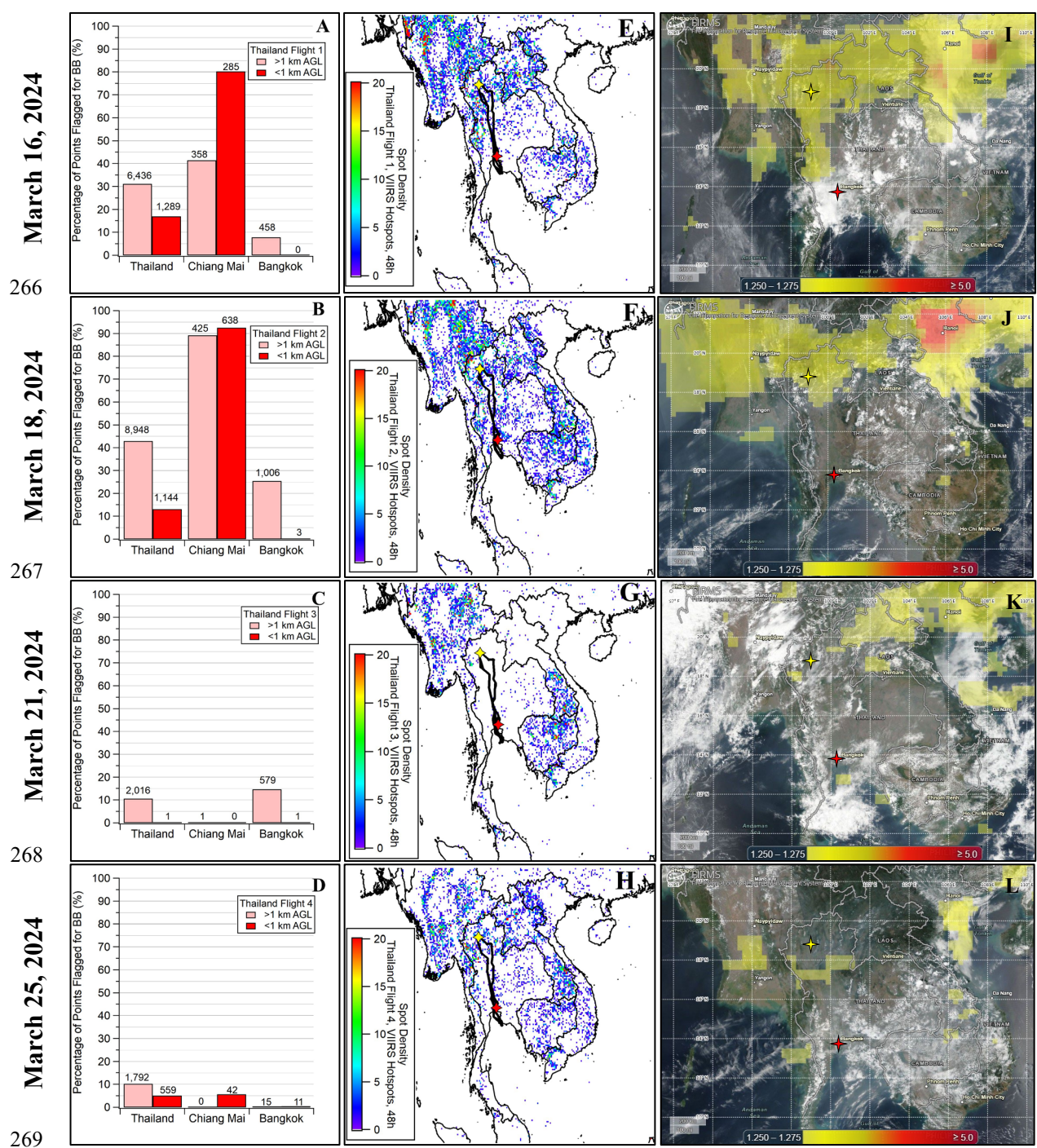


Figure 5. (A-D) Percentage of points flagged for BB for each Thailand flight across the whole flight track, over Chiang Mai, and over Bangkok for above and below 1 km AGL with total number of points flagged displayed, also shown in Table S5. (E-H) 48h VIIRS fire hotspot density maps over Southeast Asia for March 16, 18, 21, and 25, 2024. The DC-8 flight track is shown in black. The yellow star represents Chiang Mai, and the red star represents Bangkok. (I-L) VIIRS Corrected Reflectance (True Color) imagery with OMPS Aerosol Index overlay over Southeast Asia for March 16, 18, 21, and 25, 2024.

275 3.2.3 Air Mass Origin Breakdown

276 To better understand the air mass history at these locations, we used HYSPLIT back trajectory modeling along the flight path
277 to determine where these air masses came from and traveled through (using criteria from Table S4). Back trajectories with an
278 altitude < 1 km were then quantified over certain areas of interest for portions of the flight track that passed over Chiang Mai
279 and Bangkok and were flagged for BB. Figure 6 shows the results for Bangkok for all of the flights above and below 1 km of
280 altitude. For the first two Bangkok flights, there was influence from the south (Gulf of Thailand and Malay Peninsula) and
281 northeast (Eastern Thailand) above 1 km altitude, which switched to the east (Cambodia and Vietnam) for flight 3. Under 1
282 km altitude for Bangkok flights 2 and 4, most of the BB was local or traveled from the south, while for flight 3 BB was
283 transported from the northeast (Eastern Thailand and South China Sea). However, it should be noted that Bangkok experienced
284 <0.5% of BB-flagged data for Flight 4 (Fig 5D) so these results are mostly likely reflective of a single plume. Fig. 7 shows the
285 HYSPLIT back trajectory quantifications for Chiang Mai. For Chiang Mai above 1 km of altitude, the majority of back
286 trajectories are either from Central Thailand or Myanmar, with flight 1 dominated by BB transported from Central Thailand
287 and flight 2 by Myanmar. Under 1 km altitude, there is a similar trend with flight 1 flagged back trajectories originating from
288 south of Chiang Mai and flight 2 and 4 from the west. These results are in agreement with satellite fire hotspot retrievals (Fig.
289 5E-H), that show the highest hotspot density around and northwest of Chiang Mai for flights 1, 2, and 4 but for flight 3 (Fig
290 5G) a clearer area around Chiang Mai. Additional areas of high fire hotspot density occurred in Cambodia, Vietnam, and Laos,
291 especially in the last three flights.

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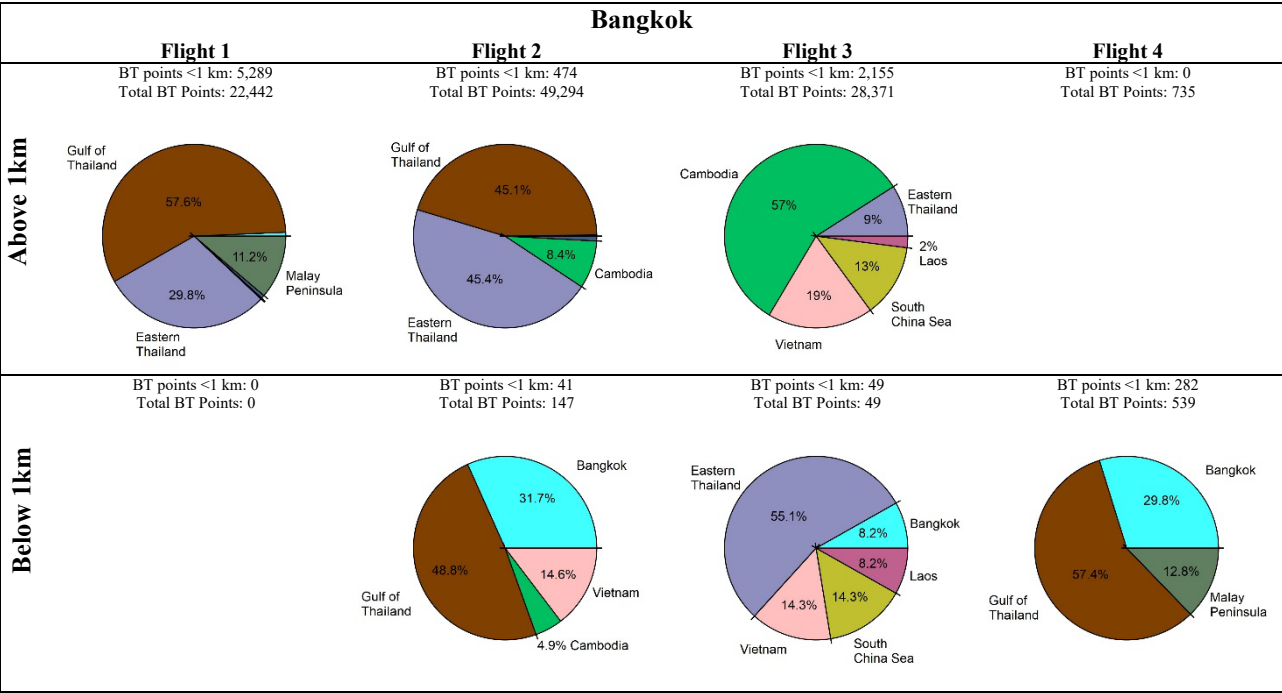


Figure 6. HYSPLIT back trajectory (BT) quantification for BB-flagged data over Bangkok for all flights above and below 1 km AGL. The number of BT points under 1 km are shown for each category along with the total number of BT points.

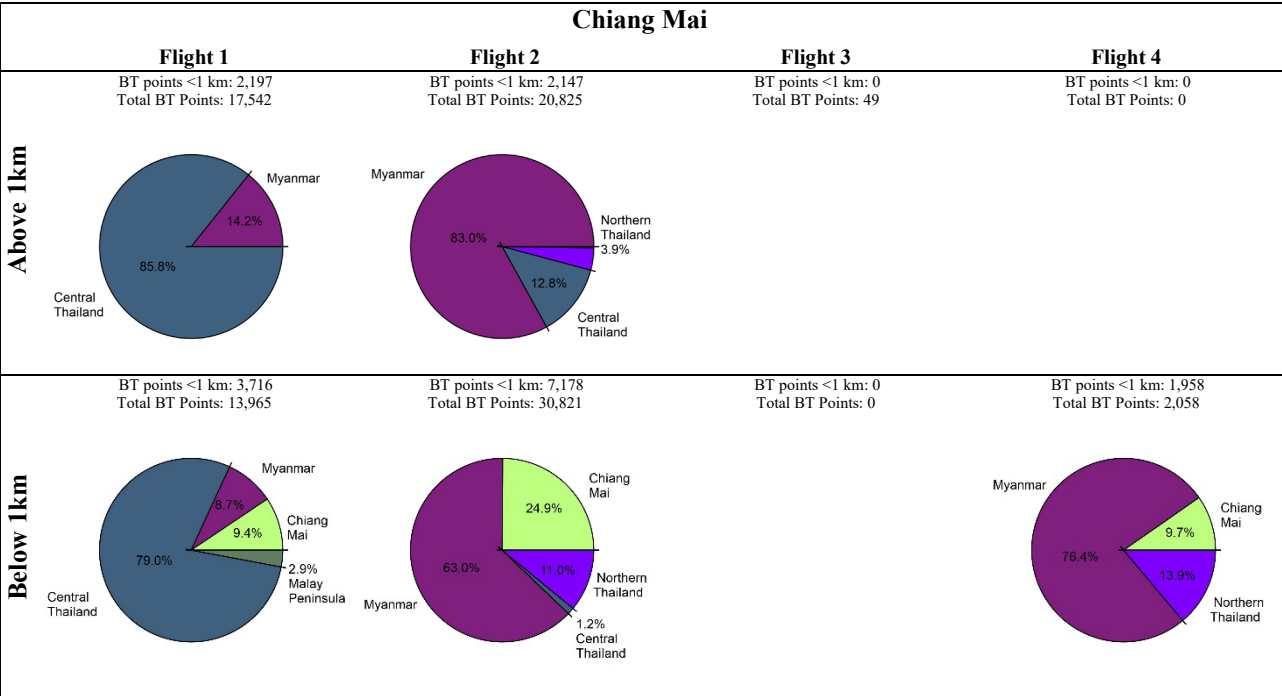


Figure 7. HYSPLIT back trajectory (BT) quantification for BB-flagged data over Chiang Mai for all flights above and below 1 km AGL. The number of BT points under 1 km are shown for each category along with the total number of BT points.

314 3.2.4 Comparison of Microphysical Aerosol Characteristics

315 To further illustrate the utility of the BB flag we have used it to compare the integrated size distribution of submicron aerosol
316 particles. Figure 8A shows particle size distributions for BB-flagged data under 1 km in Chiang Mai and unflagged data under
317 1 km in Bangkok for the same day (flight 2). The BB-flagged data has a unimodal size distribution with a mean particle
318 diameter of 150 nm (accumulation mode) compared to a bimodal size distribution in the unflagged data with a mean particle
319 diameter 27 nm (nuclei mode). The second and less significant peak in the unflagged accumulation mode (mean diameter 136
320 nm) data may be the result of background BB presence, demonstrating that this flag is optimized for clear biomass burning
321 influence and may overlook less obvious BB-influence. Several studies have demonstrated that aerosol size distributions in
322 fresh (<1 h) smoke will start at median diameters ranging from 40-150 nm and grow to larger sizes with a decrease in modal
323 width as they age (Hodshire et al., 2021; Janhäll et al., 2010; Reid et al., 1998). Figure 8B demonstrates a comparison between
324 the same location (Chiang Mai <1 km) but on flights with differing amounts of BB-influence, i.e. flights 2 and 3. While the
325 flight 3 unflagged data still exhibits smaller particle diameters (mean particle diameter 41 nm) than the flight 2 BB-flagged
326 data (mean particle diameter 150 nm), it demonstrates a relatively stronger bimodal distribution compared to flight 2 unflagged
327 data, with a mean particle diameter of 132 nm in the accumulation mode. Even though the presence of this accumulation mode
328 peak and BB markers (CO, HCN, and CH₃CN) within 20% of their respective thresholds indicate an influence from BB, the
329 dominant nuclei mode peak and $\Delta\text{CO}/\Delta\text{CO}_2$ (0.38%) and $\Delta\text{CH}_4/\Delta\text{CO}$ (104%) slopes point to urban combustion sources as the
330 dominant influence for this unflagged data.

331 When looking at black carbon mass concentrations, shown in Fig 9A, the BB-flagged data has consistently higher BC
332 concentrations compared to unflagged data across Thailand for flights 1 and 2. Flights 3 and 4 had the lowest BB-flagged BC
333 concentrations and were higher compared to the unflagged data but were within one standard deviation. This is consistent with
334 Fig 5C and 5D where flights 3 and 4 had the smallest number of points flagged for BB. The unflagged BC concentrations for
335 flights 1, 2, and 4 are also higher than unflagged flight 3 (but within one standard deviation); therefore, assuming other BC
336 emissions are consistent between the flight days, this is further evidence that the BB flag does not capture all BB-influenced
337 data, specifically BB-influence in the background. Figure 9B shows that the non-volatile number fraction of fine particles for
338 BB-flagged data is enhanced, with a narrower range compared to unflagged data across all four flights. One explanation may
339 be that the increased BC mass concentration is elevating the non-volatile number concentration of BB-flagged data in
340 combination with a decrease in the number concentration of volatile fine particles due to evaporation during transport. Future
341 work examining the age of the smoke using short-lived BB tracers could provide more information on the concentration of
342 volatile fine particles in fresh and aged smoke plumes.

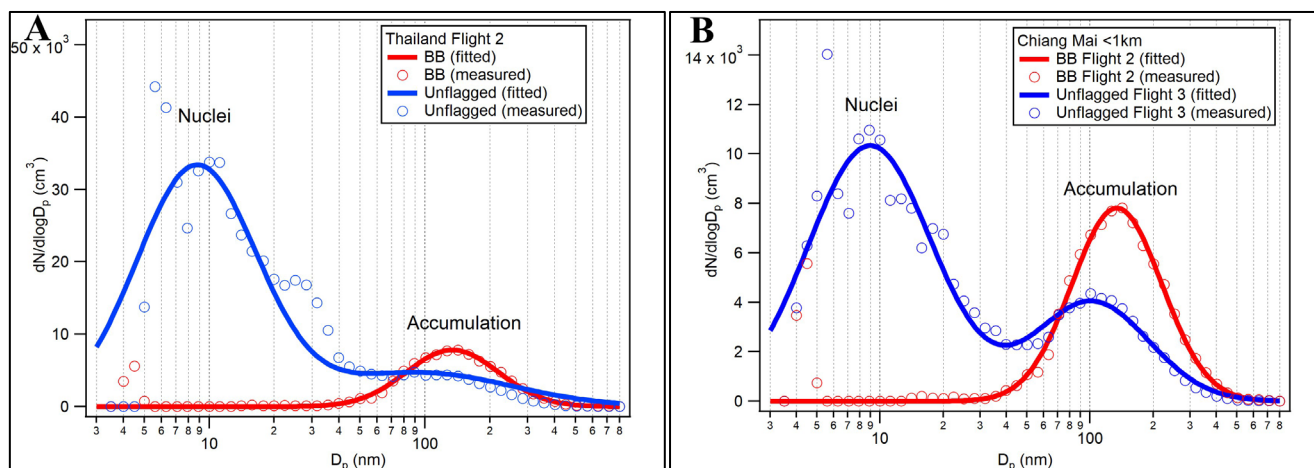


Figure 8. Fitted and measured particle size distributions for BB-flagged and unflagged conditions for flight 2 (A) and for Chiang Mai (B) < 1 km AGL.

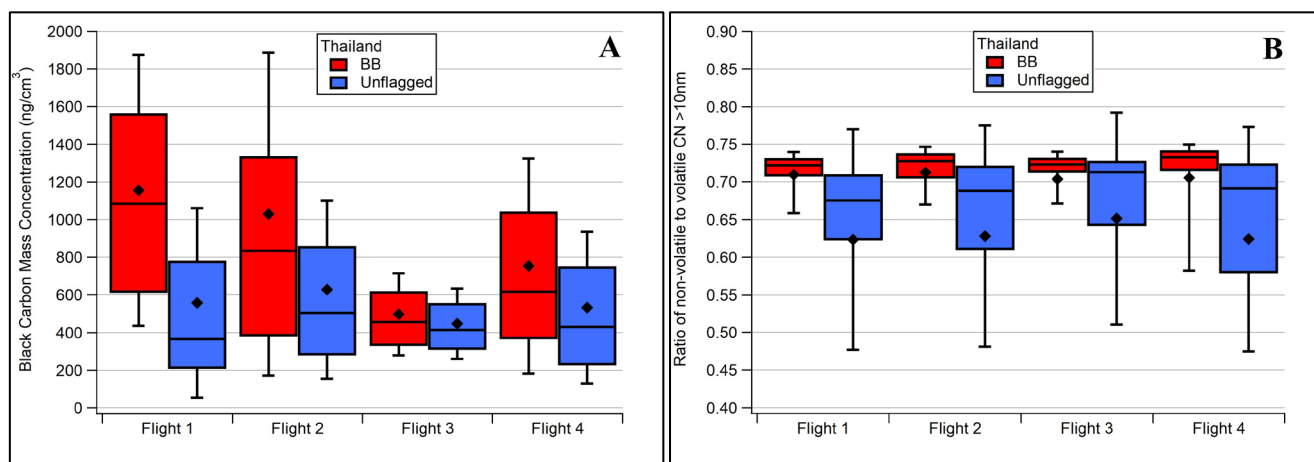


Figure 9. Black carbon mass concentrations (A) and non-volatile number fraction $CN > 10$ nm (B) for BB-flagged and unflagged data for each Thailand flight. Whiskers are representative of one standard deviation, while the boxes represent the interquartile range, the horizontal line the median, and the diamond the mean.

4 Conclusions

Biomass burning and inefficient combustion contribute to poor air quality and greenhouse gas emissions across the globe. Airborne investigations in regions prone to these emissions provide more detailed and focused measurements than typical ground or satellite methods. This work demonstrated a novel approach for identifying biomass burning-impacted airmasses in an airborne dataset. A combination of biomass burning tracers and indicators were used to distinguish biomass burning-impacted airmasses along each given flight track. The Thailand case study demonstrates the efficacy of this flag in determining areas most influenced by biomass burning and in combination with trajectory models, an idea of air mass history. Additionally, a preliminary analysis of physical characteristics of aerosols revealed differences between BB-flagged and unflagged

airmasses, including aerosol size distributions, black carbon concentrations, and non-volatile number fraction. These findings, while applicable to Asia, demonstrate the value of the method, which can be applied to other field campaigns with similar measurements. The utility of the BB flag can be increased in the future with the use of specific volatile organic compounds (VOCs) to provide information on the age of the smoke, e.g., primary BB VOCs versus secondary BB VOCs (Liang et al., 2022). Additional development of a boundary layer flag for the ASIA-AQ dataset will improve future work studying the health impacts of biomass burning smoke and inefficient combustion at ground level and transport in aloft layers.

Data Availability

All data used in this publication are open access and can be found in the ASIA-AQ Data Archive (<https://doi.org/10.5067/SUBORBITAL/ASIA-AQ/DATA001>), the FIRMS Archive (<https://firms.modaps.eosdis.nasa.gov/download/>), and the HYSPLIT model is available at <https://www.ready.noaa.gov/HYSPLIT.php>.

Author Contribution

JAM, JPD, GSD, YC, RHM, LDZ, FG, CEJ, MS, EBW, ELW, SR, YRL, KB, JDC, PW, FP, SS, WW, and AW participated in the data collection. JAM, JPD, GSD, MS, FG, KB, and WW performed the data analysis and provided feedback. JAM prepared the manuscript with contributions from all coauthors.

Competing Interests

The contact author has declared that none of the authors have any competing interests.

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392 Earth Science Data and Information System (ESDIS).

393 **References**

394 Akagi, S. K., Yokelson, R. J., Wiedinmyer, C., Alvarado, M. J., Reid, J. S., Karl, T., Crounse, J. D., and Wennberg, P. O.:
395 Emission factors for open and domestic biomass burning for use in atmospheric models, *Atmospheric Chemistry and Physics*,
396 11, 4039–4072, <https://doi.org/10.5194/acp-11-4039-2011>, 2011.

397 Anderson, T. L. and Ogren, J. A.: Determining Aerosol Radiative Properties Using the TSI 3563 Integrating Nephelometer,
398 *Aerosol Science and Technology*, 29, 57–69, <https://doi.org/10.1080/02786829808965551>, 1998.

399 Andreae, M. O. and Merlet, P.: Emission of trace gases and aerosols from biomass burning, *Global Biogeochemical Cycles*,
400 15, 955–966, <https://doi.org/10.1029/2000GB001382>, 2001.

401 ASIA-AQ Science Team: ASIA-AQ Field Campaign Data, <https://doi.org/10.5067/SUBORBITAL/ASIA-AQ/DATA001>,
402 2024.

403 Cambaliza, M. O. L., Shepson, P. B., Bogner, J., Caulton, D. R., Stirn, B., Sweeney, C., Montzka, S. A., Gurney, K. R.,
404 Spokas, K., Salmon, O. E., Lavoie, T. N., Hendricks, A., Mays, K., Turnbull, J., Miller, B. R., Lauvaux, T., Davis, K., Karion,
405 A., Moser, B., Miller, C., Obermeyer, C., Whetstone, J., Prasad, K., Miles, N., and Richardson, S.: Quantification and source
406 apportionment of the methane emission flux from the city of Indianapolis, *Elementa: Science of the Anthropocene*, 3, 000037,
407 <https://doi.org/10.12952/journal.elementa.000037>, 2015.

408 Cantrell, C. A.: Technical Note: Review of methods for linear least-squares fitting of data and application to atmospheric
409 chemistry problems, *Atmospheric Chemistry and Physics*, 8, 5477–5487, <https://doi.org/10.5194/acp-8-5477-2008>, 2008.

410 Crounse, J. D., McKinney, K. A., Kwan, A. J., and Wennberg, P. O.: Measurement of Gas-Phase Hydroperoxides by Chemical
411 Ionization Mass Spectrometry, *Analytical Chemistry*, 78, 6726–6732, <https://doi.org/10.1021/ac0604235>, 2006.

412 DiGangi, J. P., Choi, Y., Nowak, J. B., Halliday, H. S., Diskin, G. S., Feng, S., Barkley, Z. R., Lauvaux, T., Pal, S., Davis, K.
413 J., Baier, B. C., and Sweeney, C.: Seasonal Variability in Local Carbon Dioxide Biomass Burning Sources Over Central and
414 Eastern US Using Airborne In Situ Enhancement Ratios, *Journal of Geophysical Research: Atmospheres*, 126,
415 e2020JD034525, <https://doi.org/10.1029/2020JD034525>, 2021.

416 DiGangi, J. P., Diskin, G. S., Yoon, S., Alvarez, S. L., Flynn, J. H., Robinson, C. E., Shook, M. A., Thornhill, K. L., Winstead,
417 E. L., Ziemba, L. D., Cambaliza, M. O. L., Simpas, J. B., Hilario, M. R. A., and Sorooshian, A.: Technical note: Apportionment
418 of Southeast Asian Biomass Burning and Urban Influence via In Situ Trace Gas Enhancement Ratios, *EGU sphere*, 2025, 1–
419 16, <https://doi.org/10.5194/egusphere-2025-1454>, 2025.

420 Draxler, R. R.: HYSPLIT_4 User's Guide. NOAA technical memorandum ERL ARL-230, 1999.

421 Draxler, R. R. and Hess, G. D.: Description of the HYSPLIT4 modeling system, NOAA technical memorandum ERL ARL-
422 224, 1997.

423 Draxler, R. R. and Hess, G. D.: An overview of the HYSPLIT_4 modelling system for trajectories, Australian meteorological
424 magazine, 47, 295–308, 1998.

425 Fiehn, A., Eckl, M., Kostinek, J., Gałkowski, M., Gerbig, C., Rothe, M., Röckmann, T., Menoud, M., Maazallahi, H., Schmidt,
426 M., Korbeň, P., Nečki, J., Stanisavljević, M., Swolkień, J., Fix, A., and Roiger, A.: Source apportionment of methane emissions
427 from the Upper Silesian Coal Basin using isotopic signatures, *Atmospheric Chemistry and Physics*, 23, 15749–15765,
428 <https://doi.org/10.5194/acp-23-15749-2023>, 2023.

429 Guha, A., Gentner, D. R., Weber, R. J., Provencal, R., and Goldstein, A. H.: Source apportionment of methane and nitrous
430 oxide in California’s San Joaquin Valley at CalNex 2010 via positive matrix factorization, *Atmospheric Chemistry and*
431 *Physics*, 15, 12043–12063, <https://doi.org/10.5194/acp-15-12043-2015>, 2015.

432 Halliday, H. S., DiGangi, J. P., Choi, Y., Diskin, G. S., Pusede, S. E., Rana, M., Nowak, J. B., Knote, C., Ren, X., He, H.,
433 Dickerson, R. R., and Li, Z.: Using Short-Term CO/CO₂ Ratios to Assess Air Mass Differences Over the Korean Peninsula
434 During KORUS-AQ, *Journal of Geophysical Research: Atmospheres*, 124, 10951–10972,
435 <https://doi.org/10.1029/2018JD029697>, 2019.

436 Helfter, C., Tremper, A. H., Halios, C. H., Kotthaus, S., Borgeggen, A., Grimmond, C. S. B., Barlow, J. F., and Nemitz, E.:
437 Spatial and temporal variability of urban fluxes of methane, carbon monoxide and carbon dioxide above London, UK,
438 *Atmospheric Chemistry and Physics*, 16, 10543–10557, <https://doi.org/10.5194/acp-16-10543-2016>, 2016.

439 Hodshire, A. L., Ramnarine, E., Akherati, A., Alvarado, M. L., Farmer, D. K., Jathar, S. H., Kreidenweis, S. M., Lonsdale, C.
440 R., Onasch, T. B., Springston, S. R., Wang, J., Wang, Y., Kleinman, L. I., Sedlacek III, A. J., and Pierce, J. R.: Dilution impacts
441 on smoke aging: evidence in Biomass Burning Observation Project (BBOP) data, *Atmospheric Chemistry and Physics*, 21,
442 6839–6855, <https://doi.org/10.5194/acp-21-6839-2021>, 2021.

443 Holzinger, R., Warneke, C., Hansel, A., Jordan, A., Lindinger, W., Scharffe, D. H., Schade, G., and Crutzen, P. J.: Biomass
444 burning as a source of formaldehyde, acetaldehyde, methanol, acetone, acetonitrile, and hydrogen cyanide, *Geophysical*
445 *Research Letters*, 26, 1161–1164, <https://doi.org/10.1029/1999GL900156>, 1999.

446 Huangfu, Y., Yuan, B., Wang, S., Wu, C., He, X., Qi, J., de Gouw, J., Warneke, C., Gilman, J. B., Wisthaler, A., Karl, T.,
447 Graus, M., Jobson, B. T., and Shao, M.: Revisiting Acetonitrile as Tracer of Biomass Burning in Anthropogenic-Influenced
448 Environments, *Geophysical Research Letters*, 48, e2020GL092322, <https://doi.org/10.1029/2020GL092322>, 2021.

449 Janhäll, S., Andreae, M. O., and Pöschl, U.: Biomass burning aerosol emissions from vegetation fires: particle number and
450 mass emission factors and size distributions, *Atmospheric Chemistry and Physics*, 10, 1427–1439, <https://doi.org/10.5194/acp-10-1427-2010>, 2010.

452 Kurokawa, J. and Ohara, T.: Long-term historical trends in air pollutant emissions in Asia: Regional Emission inventory in
453 ASia (REAS) version 3, *Atmospheric Chemistry and Physics*, 20, 12761–12793, <https://doi.org/10.5194/acp-20-12761-2020>,
454 2020.

455 Kurokawa, J., Ohara, T., Morikawa, T., Hanayama, S., Janssens-Maenhout, G., Fukui, T., Kawashima, K., and Akimoto, H.:
456 Emissions of air pollutants and greenhouse gases over Asian regions during 2000–2008: Regional Emission inventory in ASia
457 (REAS) version 2, *Atmospheric Chemistry and Physics*, 13, 11019–11058, <https://doi.org/10.5194/acp-13-11019-2013>, 2013.

458 Liang, Y., Weber, R. J., Misztal, P. K., Jen, C. N., and Goldstein, A. H.: Aging of Volatile Organic Compounds in October
 459 2017 Northern California Wildfire Plumes, *Environ. Sci. Technol.*, 56, 1557–1567, <https://doi.org/10.1021/acs.est.1c05684>,
 460 2022.

461 Lin, G. and Wolfe, R.: VIIRS/NPP Imagery Resolution Terrain-Corrected Geolocation L1 6-Min Swath 375m NRT,
 462 https://doi.org/10.5067/VIIRS/VNP03IMG_NRT.002, 2022a.

463 Lin, G. and Wolfe, R.: VIIRS/NPP Moderate Resolution Terrain-Corrected Geolocation L1 6-Min Swath 750m NRT,
 464 https://doi.org/10.5067/VIIRS/VNP03MOD_NRT.002, 2022b.

465 Lin, N.-H., Tsay, S.-C., Maring, H. B., Yen, M.-C., Sheu, G.-R., Wang, S.-H., Chi, K. H., Chuang, M.-T., Ou-Yang, C.-F., Fu,
 466 J. S., Reid, J. S., Lee, C.-T., Wang, L.-C., Wang, J.-L., Hsu, C. N., Sayer, A. M., Holben, B. N., Chu, Y.-C., Nguyen, X. A.,
 467 Sopajaree, K., Chen, S.-J., Cheng, M.-T., Tsuang, B.-J., Tsai, C.-J., Peng, C.-M., Schnell, R. C., Conway, T., Chang, C.-T.,
 468 Lin, K.-S., Tsai, Y. I., Lee, W.-J., Chang, S.-C., Liu, J.-J., Chiang, W.-L., Huang, S.-J., Lin, T.-H., and Liu, G.-R.: An overview
 469 of regional experiments on biomass burning aerosols and related pollutants in Southeast Asia: From BASE-ASIA and the
 470 Dongsha Experiment to 7-SEAS, *Atmospheric Environment*, 78, 1–19, <https://doi.org/10.1016/j.atmosenv.2013.04.066>, 2013.

471 Lobert, J. M., Scharffe, D. H., Hao, W. M., and Crutzen, P. J.: Importance of biomass burning in the atmospheric budgets of
 472 nitrogen-containing gases, *Nature*, 346, 552–554, <https://doi.org/10.1038/346552a0>, 1990.

473 Müller, M., Anderson, B. E., Beyersdorf, A. J., Crawford, J. H., Diskin, G. S., Eichler, P., Fried, A., Keutsch, F. N., Mikoviny,
 474 T., Thornhill, K. L., Walega, J. G., Weinheimer, A. J., Yang, M., Yokelson, R. J., and Wisthaler, A.: In situ measurements and
 475 modeling of reactive trace gases in a small biomass burning plume, *Atmospheric Chemistry and Physics*, 16, 3813–3824,
 476 <https://doi.org/10.5194/acp-16-3813-2016>, 2016.

477 Nara, H., Tanimoto, H., Tohjima, Y., Mukai, H., Nojiri, Y., and Machida, T.: Emission factors of CO₂, CO and CH₄ from
 478 Sumatran peatland fires in 2013 based on shipboard measurements, *Tellus B: Chemical and Physical Meteorology*, 69,
 479 1399047, <https://doi.org/10.1080/16000889.2017.1399047>, 2017.

480 NASA FIRMS: NRT VIIRS 375 m Active Fire product VNP14IMG_T, https://doi.org/10.5067/FIRMS/VIIRS/VNP14IMG_T_NRT.002, 2024.

482 NASA VIIRS Characterization Support Team (VCST)/ MODIS Adaptive Processing System (MODAPS): VIIRS/NPP
 483 Imagery Resolution 6-Min L1B Swath 375m NRT, https://doi.org/10.5067/VIIRS/VNP02IMG_NRT.002, 2022a.

484 NASA VIIRS Characterization Support Team (VCST)/ MODIS Adaptive Processing System (MODAPS): VIIRS/NPP
 485 Moderate Resolution Bands L1B 6-Min Swath 750m NRT, https://doi.org/10.5067/VIIRS/VNP02MOD_NRT.002, 2022b.

486 Reid, J. S., Hobbs, P. V., Ferek, R. J., Blake, D. R., Martins, J. V., Dunlap, M. R., and Lioussé, C.: Physical, chemical, and
 487 optical properties of regional hazes dominated by smoke in Brazil, *Journal of Geophysical Research: Atmospheres*, 103,
 488 32059–32080, <https://doi.org/10.1029/98JD00458>, 1998.

489 Reid, J. S., Maring, H. B., Narisma, G. T., van den Heever, S., Di Girolamo, L., Ferrare, R., Lawson, P., Mace, G. G., Simpas,
 490 J. B., Tanelli, S., Ziemba, L., van Dierenhoven, B., Brientjes, R., Bucholtz, A., Cairns, B., Cambaliza, M. O., Chen, G., Diskin,
 491 G. S., Flynn, J. H., Hostetler, C. A., Holz, R. E., Lang, T. J., Schmidt, K. S., Smith, G., Sorooshian, A., Thompson, E. J.,
 492 Thornhill, K. L., Trepte, C., Wang, J., Woods, S., Yoon, S., Alexandrov, M., Alvarez, S., Amiot, C. G., Bennett, J. R., Brooks,
 493 M., Burton, S. P., Cayan, E., Chen, H., Collow, A., Crosbie, E., DaSilva, A., DiGangi, J. P., Flagg, D. D., Freeman, S. W.,
 494 Fu, D., Fukada, E., Hilario, M. R. A., Hong, Y., Hristova-Veleva, S. M., Kuehn, R., Kowch, R. S., Leung, G. R., Loveridge,
 495 J., Meyer, K., Miller, R. M., Montes, M. J., Moum, J. N., Nenes, A., Nesbitt, S. W., Norgren, M., Nowotnick, E. P., Rauber,
 496 R. M., Reid, E. A., Rutledge, S., Schlosser, J. S., Sekiyama, T. T., Shook, M. A., Sokolowsky, G. A., Stamnes, S. A., Tanaka,

497 T. Y., Wasilewski, A., Xian, P., Xiao, Q., Xu, Z., and Zavaleta, J.: The Coupling Between Tropical Meteorology, Aerosol
498 Lifecycle, Convection, and Radiation during the Cloud, Aerosol and Monsoon Processes Philippines Experiment (CAMP2Ex),
499 Bulletin of the American Meteorological Society, 104, E1179–E1205, <https://doi.org/10.1175/BAMS-D-21-0285.1>, 2023.

500 Reinecke, T., Leiminger, M., Jordan, A., Wisthaler, A., and Müller, M.: Ultrahigh Sensitivity PTR-MS Instrument with a Well-
501 Defined Ion Chemistry, *Analytical Chemistry*, 95, 11879–11884, <https://doi.org/10.1021/acs.analchem.3c02669>, 2023.

502 Sachse, G. W., Hill, G. F., Wade, L. O., and Perry, M. G.: Fast-response, high-precision carbon monoxide sensor using a
503 tunable diode laser absorption technique, *Journal of Geophysical Research: Atmospheres*, 92, 2071–2081,
504 <https://doi.org/10.1029/JD092iD02p02071>, 1987.

505 Sachse, G. W., Jr., J. E. C., Hill, G. F., Wade, L. O., Burney, L. G., and Ritter, J. A.: Airborne tunable diode laser sensor for
506 high-precision concentration and flux measurements of carbon monoxide and methane, in: *Proc.SPIE*, 157–166,
507 <https://doi.org/10.1117/12.46162>, 1991.

508 Schroeder, W., Oliva, P., Giglio, L., and Csiszar, I. A.: The New VIIRS 375m active fire detection data product: Algorithm
509 description and initial assessment, *Remote Sensing of Environment*, 143, 85–96, <https://doi.org/10.1016/j.rse.2013.12.008>,
510 2014.

511 Stein, A. F., Draxler, R. R., Rolph, G. D., Stunder, B. J. B., Cohen, M. D., and Ngan, F.: NOAA’s HYSPLIT Atmospheric
512 Transport and Dispersion Modeling System, *Bulletin of the American Meteorological Society*, 96, 2059–2077,
513 <https://doi.org/10.1175/BAMS-D-14-00110.1>, 2015.

514 Torres, O. and Goddard Earth Sciences Data and Information Services Center (GES DISC): OMPS-NPP L2 NM Aerosol Index
515 swath orbital V2, <https://doi.org/10.5067/40L92G8144IV>, 2019.

516 Tsay, S.-C., Hsu, N. C., Lau, W. K.-M., Li, C., Gabriel, P. M., Ji, Q., Holben, B. N., Judd Welton, E., Nguyen, A. X., Janjai,
517 S., Lin, N.-H., Reid, J. S., Boonjawat, J., Howell, S. G., Huebert, B. J., Fu, J. S., Hansell, R. A., Sayer, A. M., Gautam, R.,
518 Wang, S.-H., Goodloe, C. S., Miko, L. R., Shu, P. K., Loftus, A. M., Huang, J., Kim, J. Y., Jeong, M.-J., and Pantina, P.: From
519 BASE-ASIA toward 7-SEAS: A satellite-surface perspective of boreal spring biomass-burning aerosols and clouds in
520 Southeast Asia, *Atmospheric Environment*, 78, 20–34, <https://doi.org/10.1016/j.atmosenv.2012.12.013>, 2013.

521 Vay, S. A., Woo, J.-H., Anderson, B. E., Thornhill, K. L., Blake, D. R., Westberg, D. J., Kiley, C. M., Avery, M. A., Sachse,
522 G. W., Streets, D. G., Tsutsumi, Y., and Nolf, S. R.: Influence of regional-scale anthropogenic emissions on CO₂ distributions
523 over the western North Pacific, *Journal of Geophysical Research: Atmospheres*, 108, <https://doi.org/10.1029/2002JD003094>,
524 2003.

525 Warneke, C., Schwarz, J. P., Dibb, J., Kalashnikova, O., Frost, G., Al-Saad, J., Brown, S. S., Brewer, Wm. A., Soja, A., Seidel,
526 F. C., Washenfelder, R. A., Wiggins, E. B., Moore, R. H., Anderson, B. E., Jordan, C., Yacovitch, T. I., Herndon, S. C., Liu,
527 S., Kuwayama, T., Jaffe, D., Johnston, N., Selimovic, V., Yokelson, R., Giles, D. M., Holben, B. N., Goloub, P., Popovici, I.,
528 Trainer, M., Kumar, A., Pierce, R. B., Fahey, D., Roberts, J., Gargulinski, E. M., Peterson, D. A., Ye, X., Thapa, L. H., Saide,
529 P. E., Fite, C. H., Holmes, C. D., Wang, S., Coggon, M. M., Decker, Z. C. J., Stockwell, C. E., Xu, L., Gkatzelis, G., Aikin,
530 K., Lefer, B., Kaspari, J., Griffin, D., Zeng, L., Weber, R., Hastings, M., Chai, J., Wolfe, G. M., Hanisco, T. F., Liao, J.,
531 Campuzano Jost, P., Guo, H., Jimenez, J. L., Crawford, J., and The FIREX-AQ Science Team: Fire Influence on Regional to
532 Global Environments and Air Quality (FIREX-AQ), *Journal of Geophysical Research: Atmospheres*, 128, e2022JD037758,
533 <https://doi.org/10.1029/2022JD037758>, 2023.

534 Yokelson, R. J., Griffith, D. W. T., and Ward, D. E.: Open-path Fourier transform infrared studies of large-scale laboratory
535 biomass fires, *Journal of Geophysical Research: Atmospheres*, 101, 21067–21080, <https://doi.org/10.1029/96JD01800>, 1996.

536 Yokelson, R. J., Crounse, J. D., DeCarlo, P. F., Karl, T., Urbanski, S., Atlas, E., Campos, T., Shinozuka, Y., Kapustin, V.,
 537 Clarke, A. D., Weinheimer, A., Knapp, D. J., Montzka, D. D., Holloway, J., Weibring, P., Flocke, F., Zheng, W., Toohey, D.,
 538 Wennberg, P. O., Wiedinmyer, C., Mauldin, L., Fried, A., Richter, D., Walega, J., Jimenez, J. L., Adachi, K., Buseck, P. R.,
 539 Hall, S. R., and Shetter, R.: Emissions from biomass burning in the Yucatan, *Atmospheric Chemistry and Physics*, 9, 5785–
 540 5812, <https://doi.org/10.5194/acp-9-5785-2009>, 2009.

541 Yokelson, R. J., Andreae, M. O., and Akagi, S. K.: Pitfalls with the use of enhancement ratios or normalized excess mixing
 542 ratios measured in plumes to characterize pollution sources and aging, *Atmospheric Measurement Techniques*, 6, 2155–2158,
 543 <https://doi.org/10.5194/amt-6-2155-2013>, 2013.

544 York, D., Evensen, N. M., Martínez, M. L., and De Basabe Delgado, J.: Unified equations for the slope, intercept, and standard
 545 errors of the best straight line, *American Journal of Physics*, 72, 367–375, <https://doi.org/10.1119/1.1632486>, 2004.

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