



The mirabilite microbiocosm in a Carpathian contact cave

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Abstract. This study examines the microbial and geochemical environment surrounding mirabilite (sodium sulfate decahydrate) deposits in Izvorul Tăușoarelor Cave, located in the Romanian Carpathians. Using a metabarcoding approach, we analysed mirabilite, sediments, dipteran insects, drip water, and moonmilk deposits to investigate the microbial communities and elemental profiles linked to mirabilite formation. Elemental analysis revealed a geochemical signature in mirabilite samples that was dominated by sodium, sulfur, and calcium. Microbial profiling revealed a unique pattern: sulfur-reducing bacteria, such as *Desulfobacterota*, were absent in mirabilite samples, whereas *Pseudomonas* dominated, suggesting an alternative sulfur cycling pathway that potentially involves sulfide biooxidation. The presence of ammonia-oxidising archaea (*Ca. Nitrocosmicus*) exclusively in the mirabilite area, and of bacteria (*Nitrococcus*), indicates a possible influence from a small bat colony, which contributes minimal ammonia that may support the microbial equilibrium required for mirabilite growth. *Actinomyces*, abundant in mirabilite, may facilitate mineral crystallisation through mycelium-like structures. We propose the term “microbiocosm” to describe the interconnected network of biotic and abiotic elements surrounding the mirabilite environment, proposing a novel framework for investigating microbial contributions to this mineral formation.

1 Introduction

Caves are restrictive habitats due to the lack of light, saturated air humidity, and relatively constant temperature. The absence of light, and consequently the absence of plants, makes caves and their associated subterranean habitats deficient in nutrients, which are also unevenly distributed over space and time (Howarth and Moldovan, 2018). However, in recent decades, autochthonous microorganisms have been proven to form the base of subterranean food chains, even in oligotrophic caves. These microorganisms can supply nutrients and essential enzymes, playing a crucial role in the adaptation and evolution of subterranean organisms.

In addition to supporting life in caves, cave microorganisms participate in the biogeochemical cycle of elements such as C, N, S, Fe, and Mn, as well as in mineral dissolution and precipitation processes (i.e., Banks et al., 2010; Koning et al., 2022; Zhu



et al., 2022; Lange-Enyedi et al., 2024; Meka et al., 2024). These processes result in the deposition of carbonate speleothems, silicates, iron and manganese oxides, sulfur compounds, and nitrates. Such processes can be passive, with microbial cells acting as nucleation centres, or active, where bacteria produce enzymes involved in mineralisation or corrosion (dissolution). A notable example in caves is the dissolution of limestone and precipitation of gypsum, which has been documented by various authors (see review in Northup and Lavoie, 2001). Microorganisms are also involved in the oxidation, reduction, and disproportionation of sulfur compounds, reactions that facilitate assimilation and produce energy. For example, *Thiobacillus* and *Desulfovibrio* decompose sulfur-containing amino acids and proteins, making sulfur available for organisms. The activity of these bacteria depends on environmental factors such as pH, temperature, and oxygen, which are crucial for sulfate mineralisation. Sulfur-oxidising bacteria and archaea catalyse processes that transform hydrogen sulfide and elemental sulfur into sulfate. Sulfate-reducing bacteria (*Desulfovibrio*, *Desulfobacter*) oxidise sulfate into the reduced state. Biondesulfurisation is driven by *Rhodococcus* and *Pseudomonas*, which can cleave sulfur-carbon bonds, thereby removing sulfur from organic compounds.

Next-generation techniques have provided an unprecedented understanding of the diversity and complexity of microorganisms in cave formations and mineralogical processes (i.e., De Mandal et al., 2015; Hershey and Barton, 2018; Zhu et al., 2021, 2022; Akacin et al., 2022; Haidău et al., 2022; Shen et al., 2022; Samanta et al., 2023).

Using metabarcoding, we propose a new approach to studying complex microbial-driven interactions in caves, integrating several living and non-living elements in the study of mirabilite, a rare and elusive cave mineral. Mirabilite is a sulfate mineral that forms monoclinic crystals of hydrated sodium sulfate ($\text{Na}_2\text{SO}_4 \cdot 10\text{H}_2\text{O}$) and usually deposits in salt lakes. mirabilite, also known as Glauber's salt (Hill, 1979), was used in traditional medicine (Tao et al., 2024) and even mined by prehistoric populations from several areas within the Mammoth Cave System (USA; White, 2017).

It is an unstable mineral that can quickly lose water molecules and form thenardite that remains unaltered in cold and shallow marine environments (Azzaro et al., 2022). Thenardite forms under warmer conditions and at temperatures below 10°C; the hydration of thenardite forms mirabilite (Marliacy et al., 2000). Therefore, cold air and humidity, as in caves, are the prerequisites of mirabilite maintenance. It has been identified in several caves around the globe, particularly in the Northern Hemisphere (White, 2019), and in lava tubes at higher altitudes (White, 2010; Mulder et al., 2023).

The hydrated sulphates of magnesium and sodium have raised attention for their use as analogues for water availability on other planets (Pulowska et al., 2021). Microbial cells and associated beta-carotene in fluid inclusions within mirabilite suggest that these may be one of the places to search for life in Mars samples (Gill et al., 2023) or evidence for groundwater (Möhlmann and Thomsen, 2011; Battler et al., 2013).

Bacteria-induced mineral precipitation (BIMP) is the process by which bacterial activity promotes mineral formation indirectly through metabolic by-products and surface interactions with ions (Bazylnski et al., 2007; Hoffmann et al., 2021). Microorganisms can play a role in the formation and transformation of mirabilite in saline and sulfate-rich environments. While mirabilite primarily precipitates through abiotic evaporation and cooling processes, microbial activity can indirectly influence its formation by altering the geochemical conditions. Microbial communities from soil or water decompose organic



65 matter and release sulfate through oxidation processes, thus contributing to the sulfate pool available for mirabilite
precipitation. Bacteria, such as *Desulfovibrio* and *Desulfobacter*, reduce sulfate (SO_4^{2-}) to sulfide (S^{2-}), thereby depleting
sulfate from the solution and affecting local chemical equilibria. This can impact mirabilite precipitation by affecting its
supersaturation in brine environments. Therefore, even if mirabilite is abiotic, microorganisms are involved in its formation
by modifying sulfate availability through metabolic activity, acting as nucleation sites for crystal growth, and creating local
70 chemical conditions to favour precipitation.

In this study, we aimed to analyse the microbiome of the mirabilite substrate in a Romanian cave (Eastern Carpathians) by
examining the diversity of bacteria in comparison to cave sediments in areas without mirabilite, as well as in water and among
the living organisms predominantly found around the mirabilite. Our unique approach is based on the concept of the
microbioccosm of this mineral, which is relatively uncommon in caves. To understand the mirabilite microbioccosm, we
75 compared several microhabitats within the cave by incorporating the microbiomes of the living organisms. The results are
primarily expressed as the presence or absence of microorganisms in each microhabitat, and why mirabilite forms only in
small spots within a single room inside the cave. The results contribute to a deeper understanding of the involvement of
microorganisms in mirabilite formation and enhance knowledge of how a unique mineral affects or inhibits the survival of
invertebrates in cave conditions, providing insights into potential analogues for other planets.

80 **2 Materials and Methods**

2.1 Cave description

Izvorul Tăușoarelor Cave (Tausoare Cave) is situated in northern Romania, within the Rodnei Mountains, at an elevation of
942 meters above sea level (Fig. 1a). The cave measures 8.65 kilometres in length and reaches a depth of 409 meters. It is
formed within a narrow strip of sandy Upper Eocene (Priabonian)-lower Oligocene (Rupelian) limestone, interbedded with
85 thin layers of black bituminous shales (Onac et al., 2019). Fine-grained pyrite is dispersed throughout the limestone and shale.
This 60 m thick karstic formation (Silvestru and Viehmann, 1982) lies above a Middle Eocene conglomerate bed that sits on
the crystalline basement of the Rodnei Mountains (Kräuter et al., 1989).

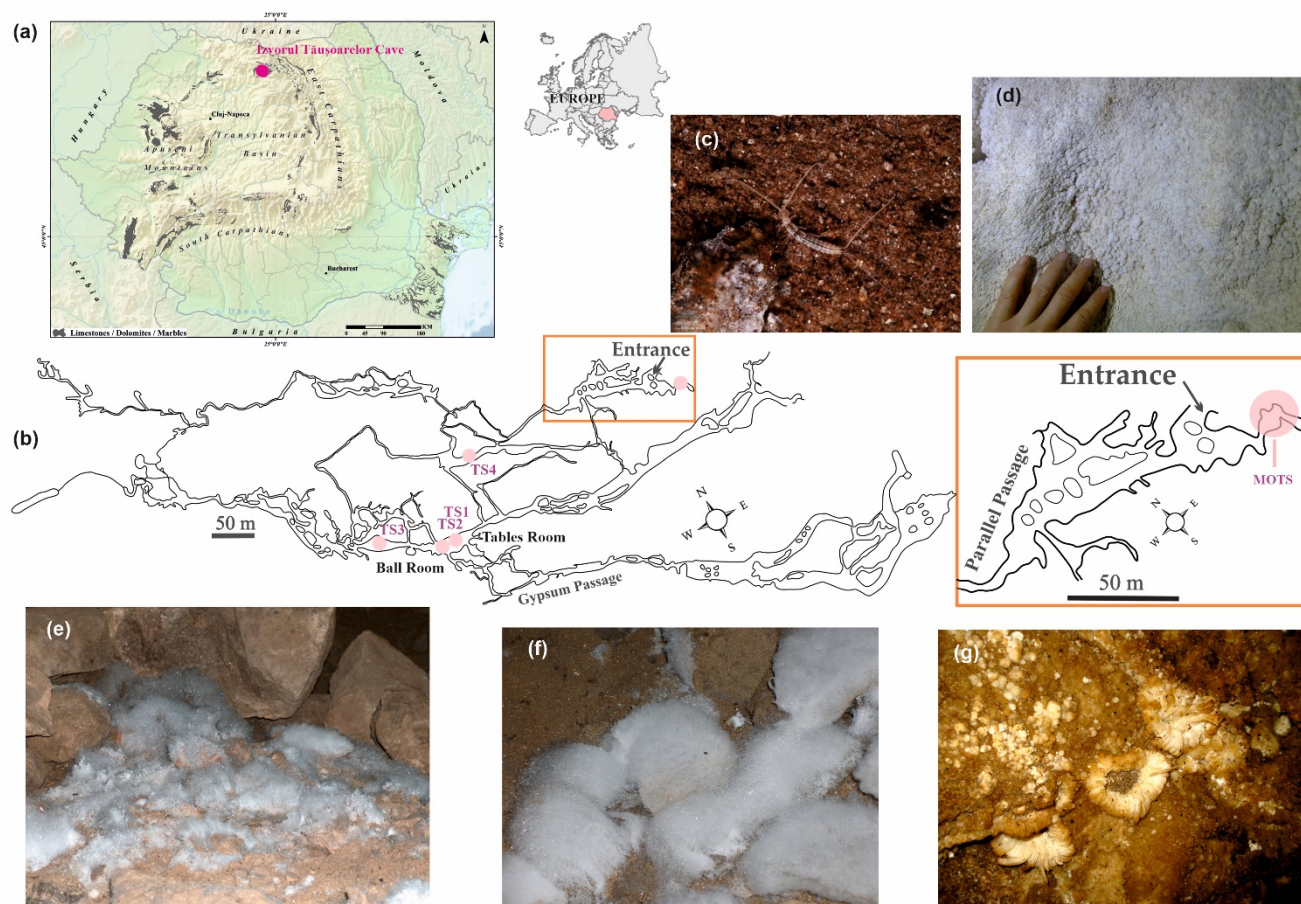


Figure 1: Izvorul Tăușoarelor Cave localisation and sampling points. a. Position of the cave in northern Romania (south-east Europe); b. Sampling stations inside the cave (see Table 1 for abbreviations); c. Endemic dipluran *Litocampa humilis comani* (Insecta); d. Moonmilk; e. Mirabilite on a sediment heap; f. Detail of the mirabilite; g. Anthodites in the Gypsum Passage. Map modified after Onac et al. (2019), photos C, D, E, and F by CTT, photo G by Traian Brad.

The water that enters the cave sinks ~300 m upstream from the entrance and flows along different cave passages, resurfacing 5.7 km from the cave terminus (Viehmann et al., 1964).

The cave's mean annual temperature ranges from 6.5 to 8°C, and humidity is consistently 100% in its deepest parts (Table 1). The cave “center” is Sala de Mese (Tables Room) from where many other passages like the Gypsum Gallery and rooms (Balls Room) continue (Fig. 1b). The Gypsum Gallery also gave the cave its fame for the unique gypsum speleothems (anthodites, crusts and crystals; Fig. 1). The primary source of sulphate ions is from the oxidation of pyrite in limestone and bituminous shale, as well as the overlying sandstone (Onac et al., 2019). Mirabilite is present in the Tables Room as patches of white, fine, needle-like crystals that can occasionally disappear (Fig. 1e-f). Its presence, combined with other minerals such as arcanite,



epsomite, syngenite, leonite, and konyaite, is related to the reaction between slightly acidic karstic water and Ca, Na, K, and Mg anions from limestone, sandstone, and clay.

The cave was declared a scientific reserve in 1965 and is currently classified with the highest level of protection afforded to a cave under Romanian legislation.

2.2 Samples description

We collected two mirabilite samples from the cave, specifically from the Tables Room (sampling stations TS1 and TS2, as shown in Fig. 1b).

One moonmilk sample was taken from a side passage near the cave entrance (MOTS, Figs. 1B and 1D). The sampling place is represented by a cave wall coated with a moonmilk layer of variable thickness (analysis of this sample is presented in Theodorescu et al., 2023). Moonmilk is a whitish, carbonatic cave deposit, with abiotic and biotic processes supposedly involved in its formation (Barton and Northup, 2007; Baskar et al., 2011; Dhimi et al., 2018). Different stages in bacteria's moonmilk formation are correlated with distinct structures and fabrics in crystal morphology (Canaveras et al., 2006).

Sediment samples were collected from the immediate vicinity of the mirabilite heap (TS1) and approximately fifty meters away (TS2), as shown in Fig. 1b, from the Ball Room (TS3), and dripping water in an active passage (TS4).

For a comprehensive view, we added the moonmilk sample and the microbiome of 10 dipluran (Insecta) individuals. The diplurans were found next to the mirabilite heap in Station 2 and in the next room, the Balls Room, in Stations 3 and 4 (Fig. 1).

The only insect adapted to cave (troglobiont) and endemic to this cave, found in relative abundance in the Tables Room, is the dipluran (Insecta) *Litocampa humilis comani* Condé, 1991 (Fig. 1c). The individuals are frequently found around the mirabilite heap, on the sediments. We collected data from 10 individuals (4 from TS2 and 6 from TS3), with TS3A and TS3B located a few meters apart.

Table 1: Analysed samples from Tausoare Cave.

Type	Station Code/Date	Temperature (°C)	Relative air humidity RH (%)	Air CO ₂ (ppm)
Mirabilite	MTS1dec	7.0	100	470
Mirabilite	MTS2dec	7.0	100	470
Sediments	STS1mar	7.0	100	470
Sediments	STS2mar	7.0	100	470
Sediments	STS1sep	7.2	100	470
Sediments	STS2sep	7.2	100	470
Sediments	STS3Asep	8.0	100	470



Sediments	STS3Bsep	8.0	100	470
Dripping water	WTS4mar	2.7	-	470
Dripping water	WTS4sep	7.1	-	470
Moonmilk	MOTSmay	7.7	90	470
Diplura microbiome	Lit_TS2mar	7.0	100	470
Diplura microbiome	Lit_TS2sep	7.2	100	470
Diplura microbiome	Lit_TS3Asep	8.0	100	470
Diplura microbiome	Lit_TS3Bsep	8.0	100	470

2.3 Physicochemical Analysis

The mirabilite, moonmilk, and sediment samples were dried in air, ground into a fine powder, and sieved through a 100-µm mesh steel sieve for the chemical analysis. The pH and electrical conductivity (EC) were measured in water and a 1/5 solid-to-water suspension using a Seven Excellence multiparameter (Mettler Toledo).

Three grams of solid samples were dissolved in 21 mL of a 3:1 (v/v) mixture of HCl (30%) and HNO₃ (65%) in a sand bath. Dripping water samples were filtered and acidified with 2 drops of HNO₃ (65%) before analysis. Na, Mg, K, Ca, Al, Fe, S, and P concentrations were measured using inductively coupled plasma optical emission spectrometry (ICP-OES) with a 5300 Optima DV (PerkinElmer, USA) spectrometer equipped with a pneumatic nebuliser. The concentrations of V, Cr, Mn, Co, Ni, Cu, Zn, Cd, Pb were measured by inductively coupled plasma mass spectrometer (ICP-MS) using an Elan DRC II (Perkin Elmer, USA) spectrometer in standard operation (Dynamic Reaction Cell in rf only mode).

2.4 DNA Extraction

Before genomic DNA extraction, cells were disrupted using FastPrep-24TM (MP Biomedicals). DNA extraction was performed using the commercial Quick-DNA Fecal/Soil Miniprep kit (Zymo Research) according to the manufacturer's protocol. Furthermore, DNA quantification was performed using the SpectraMax QuickDrop (Molecular Devices). Triplicate sub-samples were analysed for each sample.

The extracted DNA was used as a template to investigate the composition of microbial communities and was sent for 16S rRNA metagenome sequencing using a commercial service provider, Macrogen Europe. PCR of the V3-V4 hypervariable regions of the bacterial and archaeal SSU rRNA gene (Herlemann et al., 2011) was performed according to Illumina's 16S amplicon-based metagenomic sequencing protocol using the universal primers 341F and 805R.

2.5 Metabarcoding Analysis

Reads with a minimum length of 250–300 nt were analysed; Cutadapt v2.9.30 was used to remove sequencing primers and reads with N characters. Furthermore, the DADA2 package (Martin, 2011), implemented in R by adapting an existing pipeline (Callahan et al., 2016a), was used to process paired-end reads from all samples, allowing for the precise differentiation between



actual biological variation and sequencing errors. After primer removal, the resulting paired ends were loaded into the DADA2 pipeline, trimmed, filtered, and merged with a minimum required overlap of 50 nt. Chimeras were removed from merged pairs. Curated ASVs (amplicon sequence variants) were taxonomically classified using SILVA 138.1 database (Callahan et al., 2016b).

From the obtained ASVs, a mean value of triplicate sub-samples was used in the further analysis.

Sequence data generated in this study have been deposited in the European Nucleotide Archive (ENA) under the accession number PRJEB63212 for moonmilk and the rest of the samples in the National Centre for Biotechnology Information (NCBI) under accession number PRJNA1259755.

2.6 Statistical Analysis

The Phyloseq package in R (Pruesse et al., 2007) was used to process community composition and statistical differences. Mean values of replicated samples were calculated to merge their values, followed by Bray distance-based hierarchical clustering and alpha diversity analysis. The tax-glom function provided by phyloseq (Pruesse et al., 2007) was used for the taxonomic agglomeration to generate counts and relative abundances at the genus, family, and phylum levels. Only 10 counts in at least one averaged sample were considered for abundance estimation for taxa merged at each taxonomic rank.

Principal Components Analysis (PCA) was used to represent the relationship between the bacterial abundances of the samples. Agglomerative Hierarchical Clustering (AHC), based on dissimilarities and Euclidean distances, was used to cluster the bacterial abundances of the samples. Multidimensional Scaling (MDS) for the samples in a 2-dimensional space based on their physicochemical similarities. For the MDS, a similarity matrix was generated using Pearson correlation as a proximity measure between the physical-chemical parameters. Stress is a goodness-of-fit measure that varies between 0 and 1, with values near 0 indicating a better fit. The analysis was done in XLSTAT 2024.4.1.

3 Results

3.1 The mirabilite bacteriocosm

In the mirabilite samples, 1 identified ASV represented *Crenarchaeota*, and 176 ASVs belonged to *Bacteria*. *Pseudomonadota* (84% and 53%) and *Actinomycetota* (10% and 12%) were dominant in both samples. It was followed by *Bacteroidota*, abundant only in station 2 (~7%). *Gemmatimonadota*, *Firmicutes*, and *Chloroflexota* followed in lower abundances in one or both stations (Fig. 2). *Gammaproteobacteria* were the most abundant class in both samples (82% and 50%), followed by *Actinobacteria* and *Bacteroidia* in Station 2, and *Acidimicrobiia* in Station 1. *Pseudomonas* was the most abundant genus in both samples (mean value 55%), followed by *Lysobacter* (14%) in Station 2 and an unassigned *Actinomarinales* in Station 1 (5.7%; Fig. 2).

Lysobacter was abundant in all sediment samples (~33%; Fig. 2) but not *Pseudomonas*. Following in abundance were the unassigned *Longimicrobia* and *Anseongella*. The five sediment samples contained 455 Bacterial ASVs and 8 Archaeal ASVs, except Station 3, where mirabilite was absent. *Pseudomonadota* was the most abundant phylum in all samples (mean value



49%), followed by *Gemmatimonadota* (~17%) and *Bacteroidia* (~10%) only in some of the samples (Fig. 2). *Actinobacteriota*, *Acidobacteriota*, and *Firmicutes* were also abundant in some of the samples. *Gammaproteobacteria* was the dominant class in all samples, with a mean abundance of 45%, followed by *Longimicrobia* and *Bacteroidia*. *Bacteroidia* were abundant only in the September samples (Fig. 2).

190 Moonmilk composition was discussed in Theodorescu et al. (2023). The 204 identified ASVs belonged to Archaea (1) and the rest to Bacteria. As shown in Fig. 2, the most abundant genus was *Pseudomonas*, followed by 10 unassigned genera. *Pseudomonadota* (65%) was by far the most abundant phylum, represented by *Gammaproteobacteria* (62%), followed by *Patescibacteria* (18%), with the best represented *Saccharimonadia* class (10%).

195 The water samples with 494 Bacteria ASVs and 5 Archaea ASVs had *Pseudomonadota* and *Bacteroidia* as the most abundant in all samples (mean values 54% and 15%, respectively), similar to abundances in sediments (Fig. 2). *Actinobacteriota*, *Firmicutes*, and an unassigned phylum followed. Like sediments, the best-represented classes in both samples were *Gammaproteobacteria* (mean value 41%), *Alphaproteobacteria*,

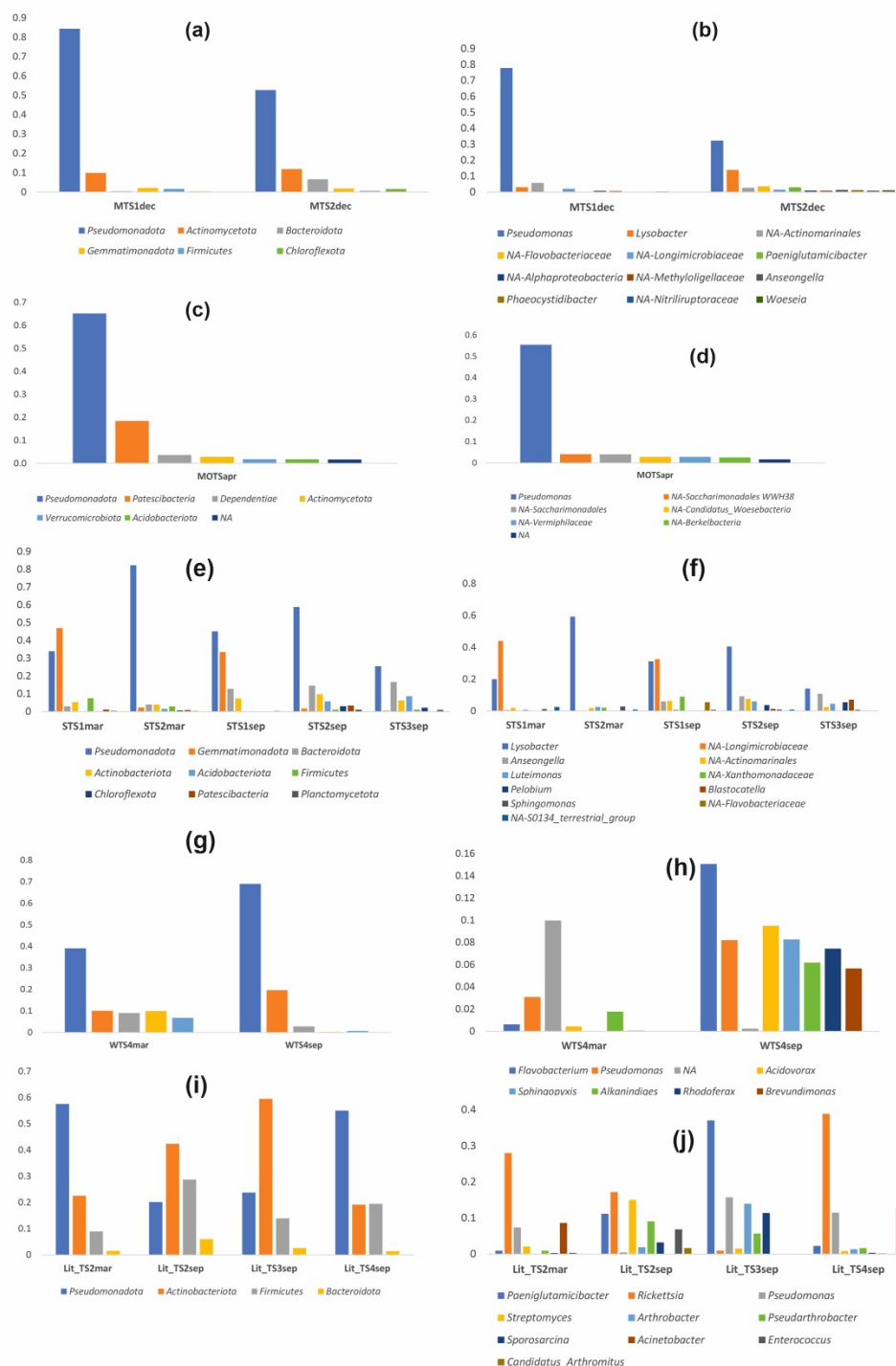


Figure 2: The most abundant bacteria in Tausoare Cave samples: mirabilite (a, b); moonmilk (c, d); Sediment (e, f); water (g, h); dipluran (i, j). See Table 1 for abbreviations.



Bacteroidia, and an unassigned class (13-20%). *Flavobacterium* was the dominant genus in the September samples, accounting for 15%, with other genera, including *Pseudomonas* and *Acidivorax*, also present. The March sample had the most abundant unassigned genus belonging to the unassigned phylum.

205 Four ASVs of Archaea and 487 ASVs of Bacteria were identified in the dipluran microbiome. The most abundant phyla were *Pseudomonadota* (20-57%), *Actinobacteriota* (19-60%), and *Firmicutes* (9-29%). *Actinobacteria* (18-59%) were the most abundant class in all the studied individuals, followed by *Alphaproteobacteria*, *Gammaproteobacteria*, *Bacilli*, and *Clostridia*, which were abundant in one or more individuals. *Rickettsia* (~20%) and *Paeniglutamibacter* (~15%) were the most abundant genera (Fig. 2).

210 Among *Archaea*, only a taxon was identified in both mirabilite samples, Ca. *Nitrocosmicus* that oxidises ammonia to nitrate, widely distributed in soils and other environments (Nicol et al., 2019). This archaeon was not identified in other samples. In sediments, only Ca. *Iainarchaeum*, *Methanobacterium*, and six other non-assigned genera were identified. Ca. *Iainarchaeum* and three non-assigned genera were identified in dipluran individuals, while water and moonmilk had only non-assigned genera.

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3.2 Bacteria involved in the sulfur cycle

Pseudomonas, which can be involved in removing sulfur from organic compounds, was the dominant genus in mirabilite (55%) and moonmilk (56%). It was also second in abundance in water (3-8%) and the dipluran microbiome (0.4-16%). Mirabilite also contained *Rhodococcus* (~0.2%). The sediments contained these genera in small abundances (<0.6%). No other sulfur

220 cycle-involved bacteria were identified in the mirabilite samples.

Sediments also contained in TS1, near the mirabilite heap, *Desulfovibrio* and *Bilophila* (each ~0.08%); both are sulphate-reducing bacteria. *Desulfobacterota* also had very low abundances (~0.003%) in TS1 and TS2. Very small amounts of other bacteria involved in the sulfur cycle were identified in moonmilk, like *Ectothiorhodospiraceae* and *Desulfobacterota* (both ~0.005%). The *Ectothiorhodospiraceae* produce sulfur globules outside their cells (Garrity, 2005), classified as slightly

225 halophilic, and *Desulfobacterota* is a sulfate-reducing phylum.

In the dipluran microbiome, one individual in TS2 had very small abundances of *Desulfovibrio* (0.02%) and other unassigned genera of *Desulfovibrionaceae*. An unassigned class of *Desulfobacterota* was present in three of the sampled individuals in very low concentrations.

230 3.3 Comparison between mirabilite and the other sample types

Pseudomonadota was dominant in mirabilite and all the other samples, with the co-dominance of *Actinomycetota* in mirabilite, *Bacteroidota* in most sediments (together with *Gemmatimonadota*) and water samples, and *Firmicutes* with *Actinobacteriota* in diplurans (Fig. 3a). The dominant class in all samples was *Gammaproteobacteria*, with one dipluran individual as an exception (Fig. 3b). In sediments, *Bacteroida* and *Longimicrobia* were the co-dominant classes, and several classes were co-



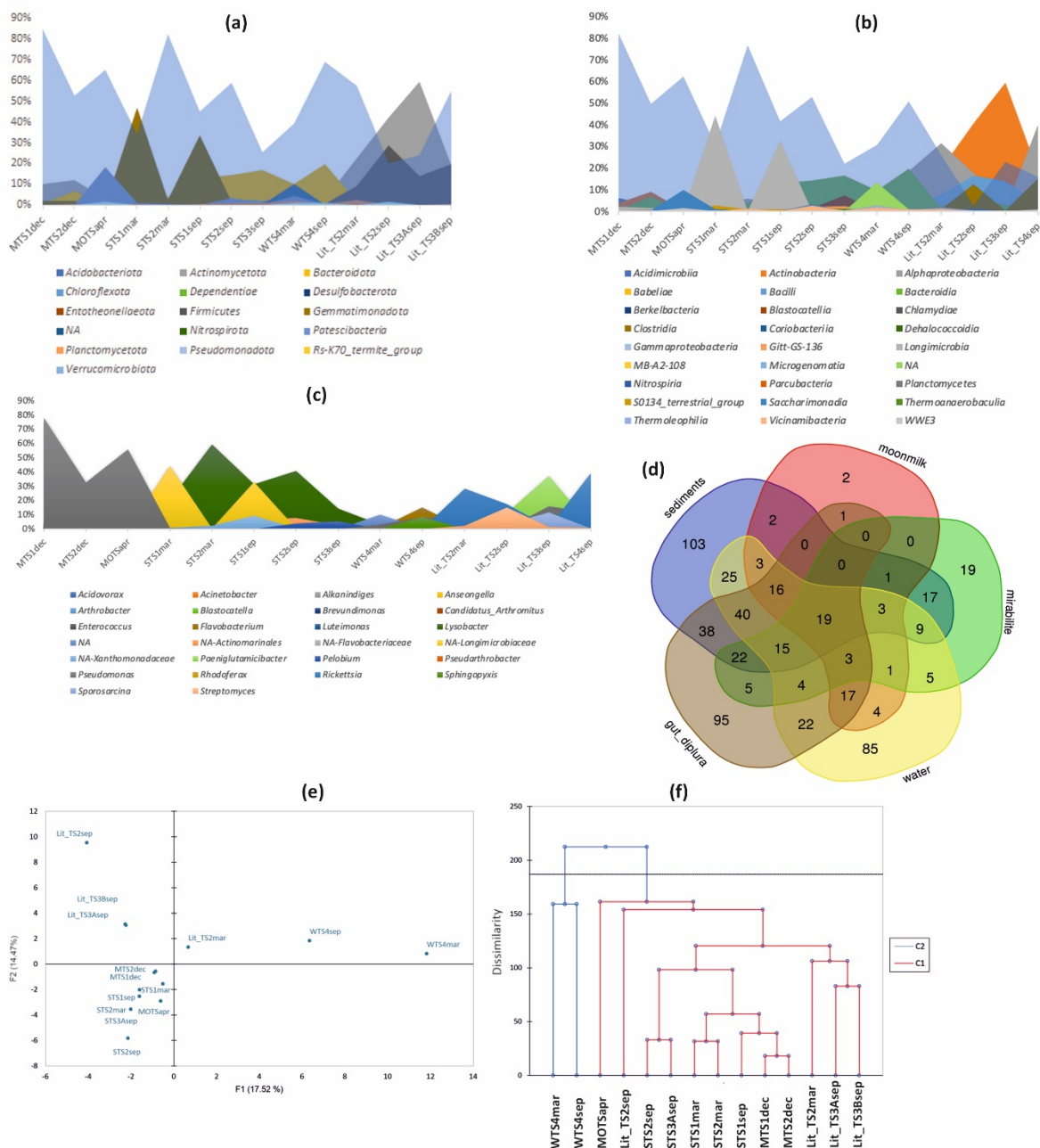
235 dominant in diplurans. Differences became most obvious at the genus level, as explained in the description above, with a clearer visualisation of them in Fig. 3c.

The Venn diagram (Fig. 3d) showed that mirabilite has 19 unique genera and another 17 common with the sediments, as mirabilite samples also included sediments from the heap. Five genera were shared with water samples, 9 with diplurans and none with moonmilk.

240 In the PCA variable space (32% total variation; Fig. 3e), the mirabilite and sediment samples are aligned along F2 axis, and so is the moonmilk sample. The dipluran individuals were separated along the same axis, while both water samples were well separated along the F1 axis from all the other samples.

The dendrogram of genus dissimilarity (Fig. 3f) more clearly shows the two main clusters, which are significantly separated, with the water samples in one cluster and all the other samples, which are dissimilar to them, at a lower significance level. The

245 moonmilk and the diplurans form separate subclusters. At the lowest dissimilarity, mirabilite clusters with a sediment in Station 1 in September are dissimilar from all the other sediment samples.



250 **Figure 3: Differences between the bacteria in the analysed samples in Tausoare Cave, with the dominant phyla (a), classes (b) and genera (c) in each sample type. d. Venn diagram of the analysed samples; e. PCA of the distribution of the most abundant bacteria genera in the variable space; f. Dendrogram showing the clustering of the samples by the most abundant bacteria genera. See Table 1 for abbreviations and Table S2 for the most abundant bacteria genera.**



3.4 Chemical comparison of the samples

Dripping water sample (WTS4) had a slightly alkaline pH (7.8 in March and 9.0 in September) and a low EC of $\sim 100 \mu\text{S/cm}$. The concentration of Ca was about 20 mg/L, whereas S was 2 mg/L, and K, Na and Mg were about 1 mg/L in both seasons. The other elements were present only in trace amounts ($<10 \mu\text{g/L}$). The moonmilk (MOTS) sample had a slightly alkaline pH (8.6), low EC ($120 \mu\text{S/cm}$) and its chemical composition was dominated by Ca (33.5%), accompanied by low amounts of Al (0.2%) and Mg, Fe, K, and P ($<0.1\%$). Sediments contain about 3-7% Ca, 0.8-1.2% Al and 0.2-2.8% Fe. The S, Na, Mg, K, and P concentrations were $<1\%$, while the trace elements (Mn, Co, Ni, Cu, Zn, and Pb) concentrations were $<0.03\%$. The sediments' pH was around 8, and EC ranged between 1400-6400 $\mu\text{S/cm}$. Generally, a higher concentration of Ca, Mg, and Fe was noticed in samples collected in September, compared to those collected in March. In contrast, the other elements were comparable in the two seasons (Table S2, Figs. 4a and S1).

There is a slight difference in the elemental composition of the two mirabilite samples compared to the other. Both mirabilite samples have an alkaline pH, but the S and Na concentrations are 5 and 2 times higher, and the Al and Fe concentrations are 5 and 15 times lower in MTS2dec than in MTS1dec. The MDS plot shows that the physicochemical parameters of the moonmilk sample are significantly different from those of the other samples (Fig. S1). However, excluding the moonmilk sample from the MDS analysis (Fig. 4b) made the differences between the water, sediments, and mirabilite evident, as well as the distinction between the two mirabilite samples.

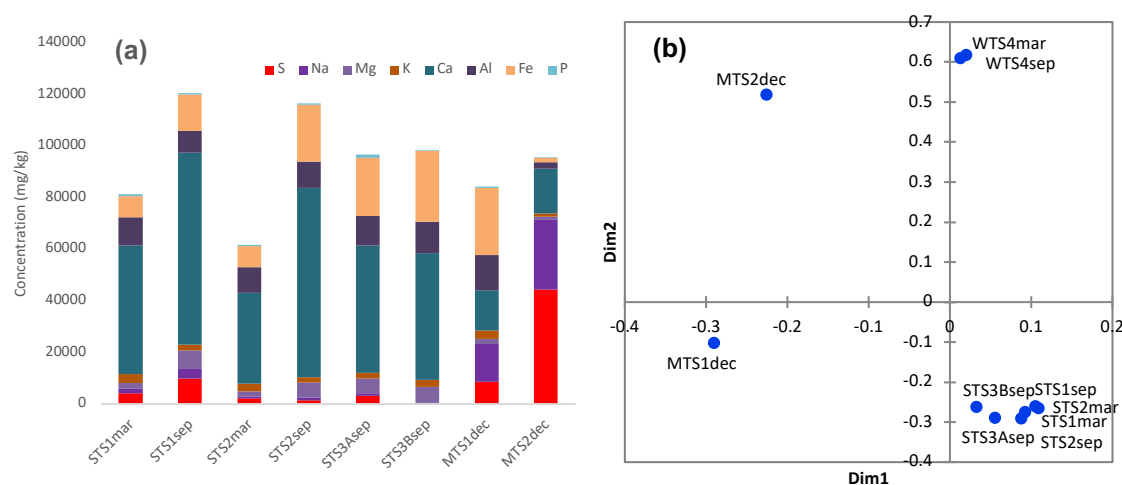


Figure 4: The chemical differences between the analysed samples from Tausoare Cave. a. Concentration of chemical elements; b. Two-dimensional MDS on the elemental concentration of mirabilite, water and sediment samples. Kruskal's stress (1) = 0.042 indicates the significance of dissimilarity between samples. See Table 1 for abbreviations and Table S2 for the physicochemical characteristics.



4 Discussion

280 We approached the metabarcoding study of mirabilite in Tausoare Cave by including all the microbiomes from and around the heap containing this rare mineral. These elements were sediments from the mirabilite area and its immediate vicinity, as well as the microbiome of dipluran insects that are typically abundant near the heaps. For comparison, we added the microbiomes of dripping water and moonmilk deposits found in other cave passages. We hypothesised that the microbial microbiome and metabolic activity associated with mirabilite could provide further insights into the conditions necessary for forming this mineral in a specific location within the cave.

285 Different elemental distribution patterns were observed across the various sample types. Sodium, calcium, and sulfur are the predominant elements in the mirabilite samples, confirming the prominence of sodium sulphate, likely accompanied by calcium carbonate and calcium sulphate. The high iron content in the MTS1dec sample accounts for secondary minerals formed by pyrite oxidation. The moonmilk completely separated from the other samples due to one order of magnitude higher Ca content. This fact aligns with the known composition of moonmilk, which has calcium carbonate as the main mineral (Theodorescu et al., 2021). Bacteria composition showed a somewhat different pattern, with water and moonmilk separating from the other samples. Mirabilite, sediment and dipluran microbiomes were more or less similar for the dominant phyla and classes, with differences at genera level.

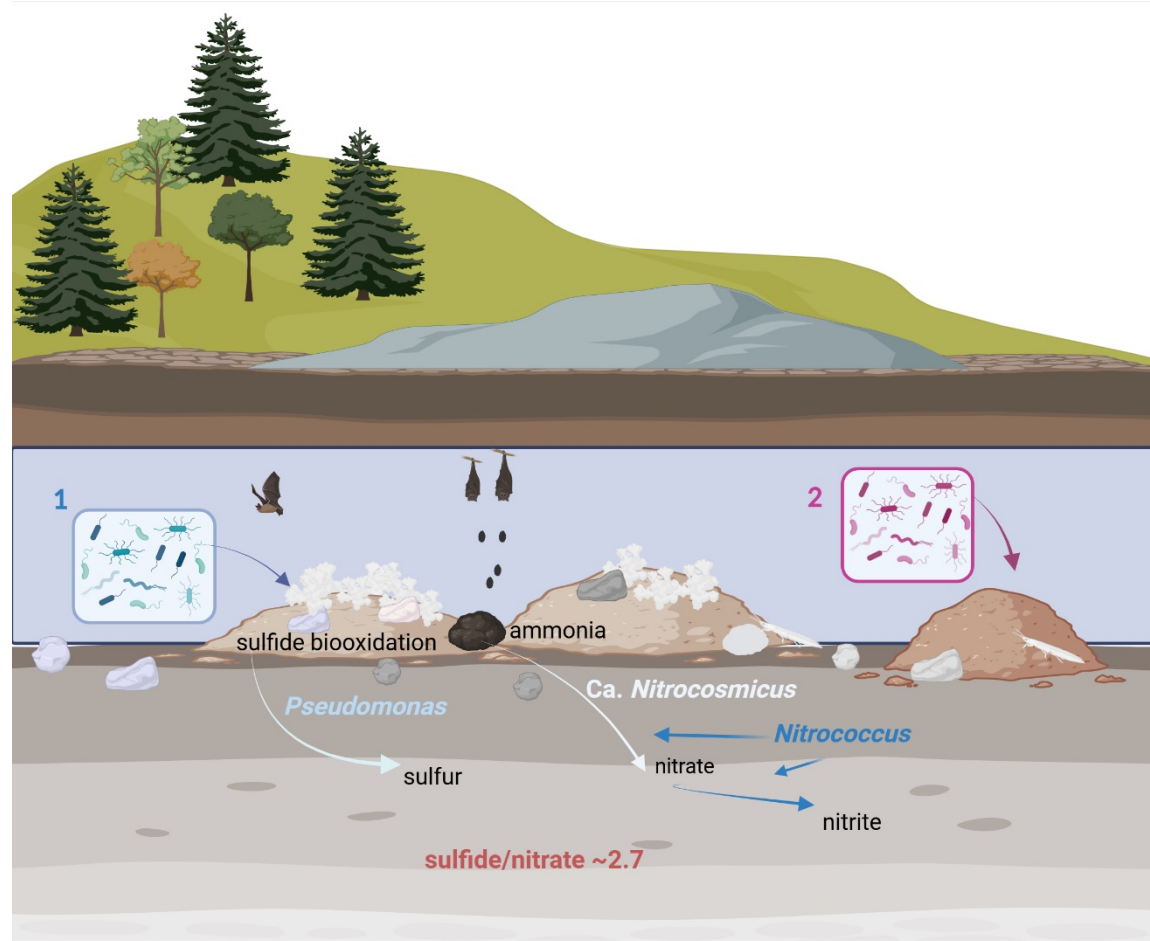
290 The first and most striking observation from analysing bacterial diversity was that no sulfur-reducing bacteria were identified in the mirabilite area of Tausoare Cave. *Desulfobacterota* and other sulfur cycle-related taxa were present in all the other sample types, although in low abundances. The absence of *Desulfobacterota* in the mirabilite samples is explained by the use of sulfur during mirabilite formation, a process sustained by the strong dominance of *Pseudomonas*. This bacterium may use organic sulfur sources, and sulfur-acquisition genes have been identified in various *Pseudomonas* representatives (see review in Scott et al., 2006). Recently, Xu et al. (2016) demonstrated that sulfide oxidation by a novel *Pseudomonas* species (C27) occurs in two stages. The first stage results in the formation of sulfur alongside nitrate reduction to nitrite, followed by the formation of thiosulfate through nitrite reduction to N₂. Biooxidation of sulfide by *Pseudomonas* sp. C27 leads to the formation of either sulfur or thiosulfate, with the ratio of sulfide to nitrate concentrations significantly influencing the oxidation to the end products. At a ratio of 0.23, the conversion of sulfide to thiosulfate increased to ~100%, but no thiosulfate was detected at a ratio of 2.7.

300 A similar pattern can be described in Tausoare Cave. The origin of sulfate in the mirabilite was proposed to be from bituminous shale, resulting from the oxidation of pyrite and marcasite within the limestone body. However, it can also have a biotic origin, such as guano deposits (Wurster et al., 2015). In the Tables Room, a small colony of bats, numbering around 50 individuals, hibernates above the mirabilite heap. This colony is too small to represent a sulfide source, but it is enough to be one of ammonia, with the ammonia-oxidising Ca. *Nitrocosmicus* (Sauder et al., 2017) only present in the mirabilite samples (Fig. 5). The interaction between *Archaea* and *Bacteria* was already emphasised in sediments of a Chinese cave (Cheng et al., 2023).



Another candidate for oxidising ammonia was *Nitrococcus*, found in the mirabilite and sediment samples, in low abundances, although it was more abundant in the mirabilite. *Nitrococcus* presence can also be explained by its role in consuming nitrites. The small amount of ammonia the bats provide is enough to sustain an equilibrium microbial process for the growth of mirabilite in this cave (Fig. 5).

315 The high abundance of *Actinomycetota*, particularly in the mirabilite samples, which can grow extensive mycelia, can serve as scaffolds for the growth of mirabilite microfibers. For this study, we have not extracted fungi.



320 **Figure 5: The schematic representation of the Tables Room in Tausoare Cave with mirabilite (1) and sediments (2) specific microbiomes.** An equilibrium between the sulfide concentration originating from the rock and the ammonia provided by bats supports microorganisms that create conditions for mirabilite formation, which is promoted by *Pseudomonas*. Created in <https://BioRender.com>.



We could not identify the possible role of dipturans in the distribution and formation of mirabilite; however, the fact that they are mainly observed around the heap area suggests that they could be a limiting factor in mirabilite dispersal, consuming some of the microorganisms that contribute to its formation. They may also rely on some sediment bacteria for nutrition, predominantly found in the Tables Room.

Mirabilite was first mentioned in the Tausoare Cave after 1972, with a large distribution between 1986 and 1987, followed by a decline and a potential recovery in the last few decades (Silvestru, 1990, and personal observations). These changes were intended to be linked to air circulation adjustments by opening new passages within the caves and subsequent modifications to air circulation and local humidity in the Tables Room. High humidity is necessary to form and maintain this mineral, which requires 10 water molecules. Disruption of air circulation can be a limiting factor in mirabilite formation. However, as the temperature deep inside the cave is rather constant and air circulation is not present in the Tables Room, we assumed this is not the case for modifications in the arrangement of mirabilite inside the cave.

Personal observations also suggest that in some periods, when tourists were more frequent in the caves (even in very low numbers), mirabilite appeared in small patches outside the classical area. It is difficult to explain this “behaviour” related to a mineral if microorganisms are not considered and their involvement is not acknowledged in creating the conditions for the formation of mirabilite. Unfortunately, there are no data on the presence, absence, or number of bats before the 21st century.

5 Conclusions

We presume that intermicrobial relationships are much more complex than those we observed by studying abundant and unique taxa in the mirabilite samples of Izvorul Tăușoarelor Cave. At this stage, we were able to provide new insights on the formation of mirabilite, as the dominance of *Pseudomonas* and the presence of an archaeon were only observed in the mirabilite samples. The relationship between the mineral kingdom and microorganisms is a domain of interest due to its numerous applications (Konhauser and Alessi, 2024); however, it remains insufficiently recognised owing to the limited number of comprehensive studies on mineral formation. Consequently, we propose using the term “microbiocosm” to examine the complex interactions of microbes, encompassing all the living and non-living microbiomes surrounding a mineral environment. This term may elucidate the processes underlying mineral formation and augment our comprehension of the intricate interactions occurring at the microscopic level.

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