



1 Long-term pig manure application increases soil organic carbon through aggregate protection 2 and Fe-carbon associations in a subtropical Red soil (Udic Ferralsols) 3 Hui Rong^{a,b}, Zhangliu Du^c, Weida Gao^a, Lixiao Ma^d, Xinhua Peng^e, Yuji Jiang^f, Demin Yan^g, Hu Zhou^{a,*} 4 ^a Key Laboratory of Arable Land Conservation (North China), Ministry of Agriculture, College of Land Science and 5 Technology, China Agricultural University, Beijing, China 6 b State Key Laboratory of Soil and Sustainable Agriculture, Institute of Soil Sciences, Chinese Academy of Sciences, 7 71 East Beijing Road, Nanjing 210008, China 8 ^c College of Resources and Environmental Sciences, China Agricultural University, Beijing 100193, China 9 ^d State Key Laboratory of Vegetation and Environmental Change, Institute of Botany, Chinses Academy of Sciences, 10 Beijing 100093, China. 11 e Institute of Agricultural Resources and Regional Planning, Chinese Academy of Agricultural Sciences, Beijing 12 100081, China 13 f College of Resources and Environment, Fijian Agriculture and Forest University, Fuzhou, 350002, China. 14 ^g Forest Fire Research Center, Nanjing Forest Police College, Nanjing 210023, China. 15 Correspondence: Hu Zhou (zhouhu@cau.edu.cn) 16 Abstract 17 Manure is known to improve soil organic carbon (SOC) in Fe-rich red soils, while the underlying 18 stabilization mechanisms remain poorly understood. In this study, four treatments were selected: (1) no amendment (Control), (2) low manure (LM, 150 kg N ha⁻¹ yr⁻¹), (3) high manure (HM, 600 kg 19 $N \; ha^{\text{--}1} \; yr^{\text{--}1}), \; (4) \; high \; manure \; with \; lime \; (HML, \; 600 \; kg \; N \; ha^{\text{--}1} \; yr^{\text{--}1} \; plus \; 3000 \; kg \; Ca \; (OH)_2 \; ha^{\text{--}1} \; 3yr^{\text{--}1}).$ 20 The quantity and quality of topsoil (0-20 cm) organic carbon were investigated by physical 21 22 fractionation, ¹³C-nuclear magnetic resonance (NMR) spectroscopy and thermogravimetry (TG) 23 analysis. Manure application increased total SOC by 65.1%-126.7% (primarily in the particulate 24 organic matter (POM) fraction), while the mineral-associated organic matter fraction (MAOM), 25 despite its higher C content (4.18-7.09 g C kg⁻¹), contributed less (65.4%-71.0%) compared to the 26 control (82.4%). POM C was stabilized via hierarchical aggregation: fresh manure inputs acted as 27 binding nuclei, increasing macroaggregates (>0.25 mm) while reducing microaggregates (0.05-0.25

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28 mm), physically isolating labile C from microbial decomposition. Concurrently, manure 29 amendments triggered Fe-mediated chemical stabilization. Elevated pH (4.8 to 5.4-7.1) enhanced 30 non-crystalline Fe oxide (Fe₀) content (+25.4%), which positively correlated with MAOM C (R² = 31 0.56, P < 0.05). Despite a chemical composition shift toward aliphaticity and reduced aromaticity, 32 thermally stable organic matters increased by 8%-12%, revealing critical role of Fe₀ (aggregates 33 were destroyed before TG analysis) in offsetting inherent molecular lability. Overall, this study 34 establishes a dual SOC stabilization framework for subtropical red soils, highlighting physical 35 protection through aggregation processes and chemical protection via Fe-carbon associations. 36 Keywords: Particulate organic matter; Mineral-associated organic matter; Nuclear magnetic 37 resonance; Thermogravimetry analysis 38 39 1. Introduction 40 Soil organic carbon (SOC), the largest carbon reservoir in the terrestrial ecosystems, plays critical 41 roles in climate mitigation and soil multifunctionality (Amelung et al., 2020; Lal, 2004). In tropical 42 and subtropical South China, red soils (Udic Ferralsols, according to Chinese Soil Taxonomy) are 43 characterized by inherently low SOC content due to intense weathering and rapid mineralization 44 (Yan et al., 2013; Zhang et al., 2013). While manure application has been widely adopted to enhance 45 SOC in these soils (Bai et al., 2023; Nichitha et al., 2023; Zhang et al., 2023), the mechanisms 46 governing SOC stabilization remain elusive. Discrepancies continue regarding whether manure 47 predominately increases SOC through chemical recalcitrance, physical protection (by aggregation 48 process), or organo-mineral interactions—given the iron-rich mineralogy of red soils and their pH-49 dependent reactivity (Kleber et al., 2021; Six et al., 2002; Song et al., 2022). 50 Existing studies present conflicting evidence on SOC stabilization pathways. For instance, 51 Mustafa et al. (2021) reported increased aromatic C (chemically recalcitrant) with manure 52 application, whereas Yan et al. (2013) observed preferential accumulation of labile O-alkyl C. This 53 paradox highlights uncertainties in how manure inputs alter SOC composition. Fe oxides contribute 54 a lot to SOC stabilization in red soils (Zhang et al., 2013). These reactive Fe phases can form stable

covalent bonds between their surface hydroxyls and organic functional groups, protecting SOC from microbial decomposition (Ruiz et al., 2024). In comparison with crystalline Fe oxides (Fe_d), non-





58 larger specific surface area (Zhang et al., 2013). The ratio of Fed and Feo is dynamically regulated by pH, which can be intentionally manipulated through manure application (Liu et al., 2020; Wang 59 60 et al., 2023). The long-term impacts of manure-induced pH shifts on Fe oxide speciation and related 61 organic carbon sequestration are still not well understood, despite the known connection between 62 pH-driven Fe oxide transformation and organic matter stabilization. Additionally, previous studies 63 often isolated chemical recalcitrance, physical protection (via aggregation), or organo-mineral interactions separately, thereby neglecting integrative assessments among these pathways. To 64 65 address this knowledge gap, an integrative approach combining physical fractionation, molecular characterization, and thermal stability analysis is essential for elucidating the coupled effects of 66 67 manure on SOC quantity and quality. 68 Physically separating soil organic matter into mineral-associated organic matter (MAOM) and 69 particulate organic matter (POM) fractions helps predicting SOC dynamics better, and clarifying 70 SOC stabilization mechanisms, with POM C physically protected in aggregated and MAOM C 71 chemically protected via organo-mineral bonding (Chenu et al., 2019; Lavallee et al., 2019; Poeplau 72 et al., 2018). While physical fractionation effectively isolates operationally defined pools (Poeplau 73 et al., 2018), it fails to reveal chemical heterogeneity of SOC (Cotrufo et al., 2019; Lavallee et al., 74 2019). Solid-state ¹³C nuclear magnetic resonance (NMR) spectroscopy addresses this gap by 75 quantifying carbon functional groups (alkyl, O-alkyl, aromatic C), yet its reliance on hydrofluoric 76 acid (HF) pretreatment risks altering native organo-mineral interactions (Kögel-Knabner, 1997). 77 Conversely, thermogravimetry (TG) provides rapid assessment of SOC thermal stability without 78 requiring pretreatment (Gao et al., 2015). 79 We hypothesized that: 1) Manure application enhance MAOM C formation by increasing non-80 crystalline Fe oxides (Fe₀), induced by elevated pH; 2) The physical protection was strengthened 81 after manure application due to the soil aggregation process, which triggered labile SOC protection; 82 3) The application of pig manure strengthened the recalcitrance of SOC, thus improved thermal 83 stability. The specific objectives of this studies were: 1) to evaluate the changes of Fe oxides and its 84 effect on MAOM formation; 2) to explore how soil aggregation affected POM formation; 3) to evaluate the effect of manure application on SOC composition and stability. 85

crystalline Fe oxides (Fe₀) exhibit organic matter adsorption capacity primarily attributed to their





86 87 2. Materials and methods 88 2.1. Site description and experimental design 89 The long-term field experiment is located at Yingtan National Agroecosystem Field Experiment 90 Station of the Chinese Academy of Sciences (28°15′20″N, 116°55′30″E) in Jiangxi Province, China. 91 The site has a typically subtropical humid monsoon climate with a mean annual temperature of 17.6°C 92 and precipitation of 1795 mm (Jiang et al., 2018). The soil is derived from Quaternary red clay, and 93 is classified as Udic Ferralsols according to Chinese Soil Taxonomy. The soil contains 36.3% clay, 94 45.2% silt and 21.2% sand. 95 The field experiment was initiated in 2002. Four treatments were compared: (1) no manure amendment (Control), (2) low pig manure with 150 kg N ha⁻¹ a⁻¹ (LM), (3) high pig manure with 96 $600 \text{ kg N ha}^{-1} \text{ a}^{-1}$ (HM), and (4) high pig manure with $600 \text{ kg N ha}^{-1} \text{ a}^{-1}$ and lime (HML). The four 97 treatments received solely pig manure as the nitrogen source, with no synthetic fertilizers applied. 98 The pig manure, collected from nearby pig farms, contained an average total carbon of 386.5 g kg 99 100 ¹, total nitrogen of 36.2 g kg⁻¹ and total phosphorus of 21.6 g kg⁻¹ on a dry matter basis. The annual 101 amount of pig manure applied to each treatment was calculated based on its nitrogen content. Since 102 all aboveground residues (stalks and leaves) and manually recoverable roots were completely 103 removed from the field after harvest, the total carbon inputs to the soil were derived exclusively 104 from pig manure. This resulted in average annual carbon inputs of 1.6 and 6.4 Mg C ha-1 for the 105 LM and HM treatments, respectively (see Supplementary Material for calculation details). The field 106 experiment was set up following a completely randomized design, with each treatment has three replicate plots. Each plot has a size of 2 m × 2 m. Lime was applied at 3 000 kg Ca (OH)2 ha⁻¹ (3a) 107 108 ¹ for the HML treatment. The field was planted with corn (Zea mays L.) monoculture annually from 109 April to July. All the management measures, including sowing, harvesting and weeding, were manually operated. 110 111 2.2 Sampling 112 Sampling was conducted in July 2019, after the harvest of corn. Triplicate topsoil samples (0-20 113 cm) were randomly collected with a shovel from each plot and composited together to form one

bulk sample. The soil samples were air-dried at room temperatureand were gently crushed with a





115 rubber mallet to pass through an 8-mm sieve, preserving aggregates >8mm for further analysis. 116 Visible plant residues, roots and stones were removed (Soil Survey Staff, 2011). 117 2.3 Soil properties measurements 118 SOC and Total nitrogen (TN) were determined by an elemental analyzer (Vario MACRO, 119 Elementar, Germany). Soil pH was measured using a glass electrode (PHS-3D, SANXIN, China) 120 with soil: deionized water ratio of 1: 2.5. Crystalline (Fed) and non-crystalline Fe oxides (Feo) were 121 extracted by DCB (Dithionite-citrate-bicarbonate) and oxalate, respectively (Yan et al., 2013), and 122 then were determined by graphite furnace atomic absorption spectrometry (GFAAS) (PinAAcle 123 900T, PerkinElmer, America). Water stability of aggregates was tested using the fast-wetting method 124 following Le Bissonnais (1996). Aggregate stability was expressed as mean weight diameter 125 (MWD). Detailed experimental processes and calculation can be found in Zhou et al. (2019). 126 2.4. Physical fractionation 127 Soil was fractionated into MAOM (<53 μm) and POM (>53 μm) fraction following Cambardella 128 and Elliott (1992) and Cotrofu et al. (2019). Briefly, 10 g sieved samples (<2 mm) were completely 129 dispersed in dilute sodium hexametaphosphate ((NaPO₃)₆, 0.5%) at a soil: solution ratio of 1:4 by 130 shaking for 18 h (25°C, 180 r min⁻¹). After dispersing, soil slurry was passed through a 53 μm sieve 131 and rinsed several times with deionized water. The fraction passing through the sieve was collected 132 as MAOM fraction, and that remaining on the sieve was collected as POM fraction. Both fractions 133 were centrifuged and the solution was decanted, and then the remaining material was oven-dried to 134 constant weights at 60°C. SOC concentration of each fraction was measured using the wet oxidation method. Dried mass proportions of each fraction (g fraction g⁻¹ soil) were calculated as follows: 135 136 $f_M = m_M / m_{bulk}$ 137 $f_P = m_p / m_{bulk}$ where f_M and f_P were the dried mass proportions of MAOM and POM fraction (g fraction g^{-1} bulk 138 soil), respectively; mm and mp (g)were the dried masses of MAOM and POM fractions; mbulk (g) 139 140 was the dried mass of bulk soil. 141 SOC in the MAOM and POM fractions were called as MAOM C and POM C, respectively in this paper. MAOM C and POM C were calculated by multiplying the dried mass proportions of each 142 143 fraction (g fraction g-1 soil) by the respective SOC concentrations (g C kg-1 fraction) as follows





144	(Garten and Wullschleger, 2000; Lian et al., 2015):
145	$MAOM C = f_M \times SOC_M $ (3)
146	$POM C = f_P \times SOC_P $ (4)
147	where f_M and f_P were calculated by Equation (1) and Equation (2); SOC $_M$ and SOC $_P$ were the SOC
148	concentrations in the MAOM and POM fraction (g C kg ⁻¹ fraction), respectively.
149	The contributions of MAOM C and POM C to total SOC (%) was calculated as:
150	Contribution of MAOM C (%) = MAOM C / Total SOC \times 100 (5)
151	Contribution of POM C (%) = POM C / Total SOC \times 100 (6)
152	where MAOM C and POM C were derived from Equations (3) and (4), respectively; total SOC was
153	the content of SOC in the bulk soil.
154	2.5. SOC chemical composition and chemical stability
155	SOC chemical composition was analyzed with a solid-state cross-polarization magic angle
156	spinning (CPMAS) 13 C nuclear magnetic resonance (NMR) spectroscopy. Prior to NMR analysis,
157	$<$ 2 mm air-dried soils were pretreated with hydrofluoric acid (HF) to remove paramagnetic Fe $^{3+}$
158	(iron oxides) following Gao et al. (2021). Firstly, 20 g soil was mixed with 100 mL 10% (w/w) HF
159	solution in a polyethylene bottle and then shaken for 0.5 h per day for three days. Afterwards, the
160	supernatant liquid was discarded and another 100 mL 10% HF solution was added again. The above
161	procedures were repeated for 15 times. The residue was rinsed 10 times with deionized water until
162	the pH was close to neutral. The remaining soil was freeze-dried and then ground in an agate mortar
163	to pass through a 100-mesh sieve (0.149 mm) for further analysis. This fine grinding ensured
164	homogeneous packing in the NMR rotor, minimizing signal heterogeneity (Simpson & Simpson,
165	2012).
166	Carbon functional groups were determined with the Bruker Ascend $500\mathrm{MHz}$ NMR spectrometer
167	(Bruker BioSpin, Rheinstetten, Germany). Dry powdered samples were placed in a 4-mm sample
168	rotor operating at a 13 C resonance frequency of 125.8 MHz. The NMR spectrometer run at a spinning
169	rate of 5kHz, and 10500 scans were collected for each sample. The spectra were collected over an
170	acquisition time of 12 ms and a recycle delay of 0.8 s. We assigned the obtained spectra to four
171	different carbon functional groups, i.e., alkyl C (0-45 ppm), O-alkyl C (45-110 ppm), aromatic C
172	(110-160 ppm) and carbonyl C (160-220 ppm) according to Kögel-Kanbner (1997). The relative

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concentrations of the different functional groups were calculated as the percentage of their peak areas to the total aeras using MestReNova 14.0 software (Mestrelab Research, 2019, Spain). Indices used to evaluate SOC recalcitrance include: alkyl C/O-alkyl C, aromaticity, aromatic C/O-alkyl C, aliphatic C/aromatic C and aliphaticity (Baldock et al., 1997; Du et al., 2017). Aromaticity = aromatic C / (alkyl C + O-alkyl C + aromatic C); Aliphatic C = alkyl C + O-alkyl C; Aliphaticity = (alkyl C + O-alkyl C) / (alkyl C + O-alkyl C + aromatic C). The alkyl C/O-alkyl C ratio reflects the degree of microbial transformation, with higher values indicating advanced decomposition and accumulation of recalcitrant alkyl compounds (Baldock et al., 1997). The aromatic C/O-alkyl C ration aligns with the alkyl C/O-alkyl C ration in reflecting the degree of SOC decomposition, whereas the aliphatic C/aromatic C ratio presents a contrary perspective to them. Aromaticity is a chemical concept denoting the resistance to microbial degradation (Kögel-Knabner, 1997). Aliphaticity is used to quantify the proportion of labile aliphatic components relative to stable aromatic moieties. 2.6. Thermogravimetry analysis TG analysis was performed using a Netzsch TG 209F1 (Netzsch-Gerätebau GmbH, Selb, Germany). Air-dried soil samples were first sieved through a 2-mm mesh, and then ground in a ball mill to pass through a 50 μ m sieve. The grounded soil samples (5 ~ 10 mg) were placed in an Al₂O₃ crucible covered with an aluminum lid and were oxidized in an atmosphere of 20 mL min-1 of synthetic air (20% O2 and 80% N2) and 20 mL min⁻¹ of N2 as a protective gas. The temperature program included a heating rate of 10°C min⁻¹ from 40°C up to 800°C. The sample mass percentage relative to the initial mass as a function of temperature was recorded simultaneously, and its first derivative (DTG) was calculated to represent the mass loss rate. Three processes were detected in the temperature range of 40°C to 800°C: hygroscopic moisture evaporation, SOM decomposition and carbonate breaking down (Gao et al., 2015; Siewert, 2004). Based on the observed DTG curve, the weight loss between 180°C and 530°C, representing the temperature range during which SOM was decomposed, was defined as the total mass of SOM (Exotot). The mass loss between 180°C and 380°C was a consequence of thermally labile SOM oxidation (Exo₁) and that between 380°C and 530°C was caused by the combustion of more stable organic matter (Exo2) (Gao et al., 2015; Volkov et al., 2020). We used two parameters, the ratio of





202 Exo1 and Exotot (Exo1/ Exotot) and the temperature at which half of the SOM was decomposed (TG-203 T₅₀), to characterize the thermal stability of SOM. Higher Exo₁/ Exo_{tot} and lower TG-T₅₀ values 204 indicate that the sample has more thermally labile or unstable SOM (Gao et al., 2015; Siewert, 2004). 205 2.7. Statistical analysis 206 Statistical analyses were carried out with R Studio software (R Development Core Team, version 207 4.1.2). One-way analysis of variance (ANOVA) was conducted to assess the effect of amendments 208 on soil physico-chemical properties, SOC physical fractions, chemical composition and thermal 209 indices. The independence of samples, normality of residues and homogeneity of variances were 210 checked by Chisq, Shapiro-Wilk and Bartlett test, respectively. Fisher's least significant difference (LSD) method was used for the multiple comparisons of means with a 0.05 significance level. Linear 211 212 regression analysis was conducted to investigate the correlations between SOC and iron oxides, soil 213 aggregation. Principal component analysis (PCA) was performed to evaluate the relationship 214 between the quantity and quality of SOC and factors related to chemical protection and physical 215 protection. Pearson correlation analyses were performed to explore the relationships between 216 chemical composition and thermal indices. 217 3. Results 218 3.1. Long-term pig manure application increased SOC, TN, pH, non-crystalline (Feo) and 219 improved soil aggregation 220 Long-term manure amendment altered the soil physic-chemical properties (Table 1). Relative to 221 Control, the LM, HM and HML treatments increased SOC concentration by 64.9%, 116.2% and 126.6%, respectively (P < 0.05), and increased TN concentration by 48.0%-108.2% (P < 0.05). The 222 223 pH values were increased by 0.62-2.28 units after pig manure application (P < 0.05). Application of 224 pig manure had no significant effect on crystalline iron oxides (Fe_d) content (P > 0.05), but the HM 225 and HML treatments significantly increased non-crystalline iron oxides (Fe_o) by 25.4% (P < 0.05). 226 Relative to Control, the HM and HML treatments significantly increased macroaggregates (>0.25 mm) content by 15.8% and 16.8%, respectively, and they increased MWD by 24.3% and 35.0%, 227 228 respectively (P < 0.05). Microaggregates (0.05-0.25 mm) was decreased by 30.4% and 36.4%, 229 respectively under the HM and HML treatments (P < 0.05). 230 3.2. Long-term pig manure application affected SOM physical fractions: MAOM and POM





231	The distribution of MAOM and POM fractions was significantly influenced by manure and lime
232	amendments ($P < 0.05$, Table 2). Across all treatments, MAOM dominated the soil mass proportion
233	(72.5%-75.1%), whereas POM mass proportion increased progressively from 16.6% in the control
234	to 19.3%–19.4% under the HM and HML treatments.
235	Manure application improved SOC concentration in both fractions, with increase rates of 32.0%-
236	66.8% and 208%-592% in the MAOM and POM fractions, respectively. Despite this,
237	the contribution of MAOM to total SOC declined from 82.4% in the control to $65.5\%-65.8\%$ under
238	the HM and HML treatments, while POM contribution increased nearly threefold (from 8.8% to
239	23.7%-26.0%). Lime addition (HML vs. HM) did not significantly alter mass proportions but
240	further enhanced POM C concentration (+15.4%) and its SOC contribution (+9.7%).
241	3.3. Effect of long-term pig manure application on SOC chemical composition and recalcitrance
242	The solid-state ${}^{13}\mathrm{C}\mathrm{NMR}$ spectra showed different signal patterns for the different treatments (Fig.
243	1) and quantified the ratios of the different SOC functional groups shown in Table 3. Relative to
244	Control, the HM and HML treatments significantly increased alkyl C by 4.6%-4.9% ($P < 0.05$),
245	while the LM treatment showed no significant change ($P > 0.05$). The O-alkyl C was significantly
246	increased by 5.5%, 2.6% and 2.2% under the LM, HM and HML treatments, respectively ($P < 0.05$).
247	Aromatic C and carbonyl C were decreased by 12.4%-13.2% and 0.9%-8.7%, respectively, in the
248	manured treatment ($P < 0.05$). While aromatic C was increased in the content after manure amend,
249	its relative proportion was decreased by 12.4%-13.2% ($P < 0.05$).
250	Relative to Control, the LM treatment significantly decreased alkyl C/O-alkyl C ratio ($P < 0.05$),
251	but the HM and HML treatments had no significant effect on the ratio ($P > 0.05$, Table3). The
252	aromaticity was decreased by 14.4%, 12.9% and 13.2% under LM, HM and HML treatments,
253	respectively ($P < 0.05$, Table3). Similarly, the aromatic C/O-alkyl C ratio was decreased by 14.7%-
254	17.7% in the manured treatment ($P < 0.05$, Table3). In contrast, the aliphatic C/aromatic C ratio and
255	aliphaticity were increased by 18.1%-20.7% and 2.44%-3.66% in the manured treatments,
256	respectively ($P < 0.05$, Table3).
257	3.4. The effect of long-term pig manure application on SOC thermal stability
258	The shapes of TG and its first derivatives (DTG) curves showed distinct weight losses rate as
259	temperature increased to above 100°C, in the order of HM>HML>LM>Control (Fig. 2). Relative to





261 of 180°C to 530°C (Exo_{tot}) by 11.1% - 17.4% (P < 0.05, Table 4), and significantly increased the mass losses during 180-380°C (Exo1) and 380-530°C (Exo2) by 14.6%-26.5% and 7.8%-13.7%, 262 263 respectively (P < 0.05, Table 4). The LM and HML treatments significantly increased the ratio of 264 $\text{Exo}_1/\text{Exo}_{\text{tot}}$ (P < 0.05), whereas the HM treatment had no significant effect (P > 0.05, Table 4). HM and HML treatments significantly decreased TG-T₅₀ by 10.7-12.0 °C (P < 0.05), while LM 265 266 treatments showed no significant difference compared to Control (P > 0.05, Table 4). 267 3.5. Relationships between SOC and factors related to the mineral protection and physical 268 protection of SOC 269 Fig. 3 showed correlations between SOC and possible variables associated with the chemical 270 protection and physical protection of SOC. SOC in the bulk soil was significantly positively correlated with MAOM C (R²=0.97, Fig. 3A). While MAOM C showed no relationship with Fed 271 $(R^2=0.04, \text{ Fig. 3B})$, it was positively correlated with Fe₀ $(R^2=0.56, \text{ Fig. 3C})$. In addition, there was 272 no correlation between MAOM C and the content of clay and silt (R^2 =0.34, Fig.3D). 273 274 SOC in the bulk soil was significantly positively correlated with POM C (R^2 =0.95, Fig. 3E). POM 275 C was significantly associated with soil aggregation, evidenced by the positive correlation with 276 macroaggregates (>0.25 mm) (R^2 =0.78, Fig. 3F) and MWD (R^2 =0.67, Fig. 3H), while negative 277 correlation with microaggregates (<0.25 mm) ($R^2=-0.71$, Fig. 3G). 278 A PCA plot diagram (Fig. 4) revealed distinct associations between SOC quantity/quality and 279 stabilization mechanisms. SOC vector aligned positively with POM C, macroaggregates and MWD, 280 but inversely with microaggregates. Chemically, SOC covaried with MAOM C and Feo, yet formed 281 an obtuse angle (>90°) with Fed. TG-T50 negatively associated with O-alkyl C and aliphaticity 282 (angles $> 90^{\circ}$), but positively linked to aromatic C and aromaticity (angles $< 90^{\circ}$). 283 4. Discussion 284 4.1. POM C and physical protection 285 Long-term pig manure application caused an increase of SOC in the bulk soil, and the increase 286 was mainly derived from continuous manure inputs (Gong et al., 2009). Although maize 287 rhizodeposition (including root exudates and sloughed-off cells) (typically <10% of net primary 288 productivity; Pausch and Kuzyakov, 2018) contributes to SOC during the growing season, its annual

Control treatment, pig manure application significantly increased the total mass losses in the range

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comparison with manure inputs, which ranged from 1.6 to 6.4 Mg C ha⁻¹ yr⁻¹ in the LM and HM treatments, respectively. Furthermore, the rigorous removal of all harvest residues (stalks and recoverable roots) excluded aboveground and root biomass as a source of soil carbon deposition. Thus, the SOC increase in the LM/HM treatments (Table 1) can be predominately attributed to the exogenous organic C supplied by pig manure, and the HML treatment can be partly influenced by lime. Following microbial decomposition and transformation, the applied manure-C was progressively partitioned into distinct SOC pools. Manure-derived carbon was preferentially allocated to the POM fraction rather than the MAOM fraction. The contribution of POM C to SOC increased nearly threefold (8.8% to 26.0%), while that of MAOM C decreased significantly from 82.4% in the Control to 65.4%-71.0% under manure application treatments (Table 2). The result was in line with Lan et al. (2022) and Li et al. (2018), who reported that increased manure substitution improved the POM C/MAOM C ratio, indicating manure was beneficial to the formation of POM C. Recent biomarker evidence demonstrated that plant-derived carbon contributes disproportionately to POM in comparison with MAOM fraction (Zou et al., 2023). Notably, ¹³C isotopic tracing study revealed that 70%-87% of residue-derived SOC was accumulated in the POM fraction at two sites in the study of Mitchell et al. (2021). These results jointly validate that fresh organic inputs are distributed primarily in the POM fraction due to direct occlusion within macroaggregates before microbial processing. Therefore, POM C can be a good indicator of SOC dynamics under field management (Álvaro-Fuentes et al., 2021; Wu et al., 2023). POM C was susceptible to decomposition due to its rapid dynamics, whereas it was simultaneously protected through physical occultation within soil aggregates. According to the classical hierarchical aggregation model (Six et al., 2000), organic carbon acts as a primary binding agent for microaggregates (<0.25mm) to form macroaggregates (>0.25mm), with POM serving as both a structural nucleus and a transient carbon reservoir (Six et al., 2000). In this study, macroaggregates were significantly increased, while microaggregates were significantly decreased after pig manure application (Table 1). These results implied that manure-derived carbon bound microaggregates to form macroaggregates, thus providing physical protection for SOC in the POM

carbon inputs is only 0.2-0.5 Mg C ha⁻¹ yr⁻¹ (Dennis et al., 2010). This contribution is negligible in





fraction (Peng et al., 2023). The observed positive correlation between POM C and both macroaggregates and aggregate MWD (Fig.3; Fig.4) further verified the critical role of soil aggregation. However, the physical mechanisms by which POM facilitates this process (e.g., microbial mediation, hydrophobic interactions, or polysaccharide bridging) require further investigation. Experimental approaches such as isotopic labelling combined with micro-scale imaging (e.g., electron microscopy or X-ray computed tomography) can visualize the spatial distribution of POM within aggregates and quantify its role in aggregate formation. Future study should pay attention to these new technologies.

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4.2. MAOM C and chemical protection

MAOM retained higher SOC concentrations (4.18-7.09 g C kg⁻¹ bulk soil) irrespective of treatments, though its contribution to total SOC decreased after manure application (Table 2). MAOM C exhibited significantly longer mean residence time (MRT) compared to POM C, with reported turnover periods of 26-40 years versus 2.4-4.3 years, respectively (Benbi et al., 2014; Garten and Wullschleger, 2000). This fundamental difference originates from the unique stabilization mechanism of MAOM C through persistent organo-mineral associations (Lavallee et al., 2019). While clay minerals are widely recognized as key MAOM stabilizers (Hemingway et al., 2019; Liang et al., 2017), our study reveals a distinct iron oxide-dominated mechanism in these Ferich red soils. The strong correlation between MAOM C and non-crystalline Fe oxides (Fe₀, R²=0.56; Figs. 3C, 4) contrasts with its weak relationship with clay + silt content (R²=0.34; Fig. 3D), highlighting the pivotal role of Feo in this red soil. This iron-mediated stabilization likely stems from the exceptionally high specific surface area of Fe₀ (~800 m² g⁻¹ in ferrihydrite; Kleber et al., 2005) and its superior capacity to form stable organo-mineral complexes (Eusterhues et al., 2005; Lehmann and Kleber, 2015). The increased pH after manure application (Table 1) created conditions favoring Feo preservation (Vithana et al., 2015), with an increase of 25.4% in the content of Feo. The increased Feo content provided abundant reactive surfaces for MAOM formation. This pH-Feo-MAOM nexus establishes a self-reinforcing stabilization mechanism: manure-derived organic ligands interact with Fe₀ to form stable complexes, simultaneously protecting both organic carbon and Fe_o from dissolution (Kleber et al., 2021).

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The positive correlation between Feo and MAOM C (Fig. 3C) suggested that non-crystalline Fe oxides played a critical role in stabilizing SOC through organo-mineral interactions. However, while our data support the association between Feo and MAOM C, the underlying mechanisms (e.g., adsorption, co-precipitation, or ligand exchange) remain speculative due to the lack of direct molecular-scale evidence. Future studies employing advanced spectroscopic techniques (e.g., synchrotron-based X-ray absorption spectroscopy or NanoSIMS) could explicitly characterize the binding forms of Fe-organic complexes, thereby validating the causal relationship between Feo and MAOM formation. 4.3. Chemical composition and thermal stability Input of new organic material could alter chemical composition of SOC and lead to the change of the molecule recalcitrance of SOC (Guo et al., 2019; Yan et al., 2013; Zhou et al., 2010; Zhang et al., 2013). In the current study, pig manure application increased the content of O-alkyl C, but declined that of aromatic C. Pig manure was rich in cellulose and lignin components, so the introduction of manure greatly increased O-alkyl C (Li et al., 2015). The considerable increase of O-alkyl C accounted for the decrease of the relative proportion of aromatic C and the decrease of aromaticity. Long-term pig manure application strengthened SOC thermal stability by improving the content of thermally stable organic matters while it decreased TG-T₅₀. Thermal analysis suggested that manure amend significantly increased SOM content, evidenced by the increase of the mass losses during 180-530 °C (Exotot) (Table 4). The result was consistent with the change of SOC content measured by conventional method (Table 1), indicating TG technology was promising for measuring SOM content (Siewert, 2004; Tokarski et al., 2018). The decrease of TG-T50 reflected more easily decomposable SOC accumulated after manure application (Gao et al., 2015; Siewert, 2004). The result consisted with that in the NMR spectroscopy, where greater O-alkyl C and higher aliphaticity were found under manure treatments (Table 3). SOC with more O-alkyl C functional groups or higher aliphaticity was less likely to resist thermochemical degradation as revealed by the negative relationship between TG-T₅₀ and aliphacity (Fig. 4; Fig. 5) (Hou et al., 2019; Lehman and Kleber, 2015), thus leading to a decrease of TG-T₅₀. In this study, the thermally labile organic

matters accounted for over half of the total organic matters, so the decrease of TG-T50 after manure

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application was just a result of the increased thermally organic matters not the decrease of thermal stability. In contrast, thermal stability should be strengthened according to the increased thermally stable organic matters after manure application. Since soil structure was destroyed due to the ground process of soil samples before TG analysis, the increase of the thermally stable organic matters was ascribed to mineral protection, where the correlations between organic carbon and Fe oxides increased the thermally resistance of OC (Ruiz et al., 2023). 5. Conclusion Manure application increased SOC quantity and improved its quality in Fe-rich red soils. SOC in the POM fraction exhibited the most pronounced response to manure inputs, while the majority of SOC was stored in the MAOM fraction. Furthermore, SOC was stabilized by distinct yet complementary mechanisms: physical protection via aggregation process, and chemical protection via Fe-organic associations induced by elevated pH. In addition, manure application increased thermally stable organic matters. However, to better understand inherent mechanisms, future work should focus on molecular-scale characterization of Fe-organic interactions using synchrotron techniques (e.g., Fe K-edge XANES/EXAFS), and in-situ visualization of POM-mediated aggregation through advanced imaging tools (e.g., SEM-TEM or μ -CT) coupled with 13 C-labelled manure to track POM dynamics within aggregates. Acknowledgments This work was financially supported by the NSFC-CAS Joint Fund Utilizing Large-scale Scientific Facilities (No. U1832188) and Chinese Universities Scientific Fund (2023RC047). We thank Yanyan Cai and Xing Xia for laboratory assistance. **Author contributions** HR conceived the experimental approach, took soil samples from the field, carried out the laboratory and data analyses, wrote the first draft of the manuscript, and contributed to subsequent drafts. ZLD helped analyze NMR spectrum and revise the manuscript. WDG contributed to interpretating TG data and revising the manuscript. LXM conducted the experiment of NMR. XHP helped design the

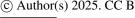
experiment, and analyze data. YJJ contributed to the design of field experiment and the process of





405 taking soil samples. DMY contributed to the process of writing. HZ contributed to funding 406 acquisition, conceiving the experimental approach, carrying out data interpretations, and the writing 407 of all subsequent manuscript drafts. 408 409 Conflict of Interest 410 Hu Zhou is a member of the editorial board of SOIL. 411 References 412 413 Álvaro-Fuentes, J., Franco-Luesma, S., Lafuente, V., Sen, P., Usón, A., Cantero-Martínez, C., Arrúe, J.L., 2021. 414 Stover management modifies soil organic carbon dynamics in the short-term under semiarid continuous 415 maize. Soil Till. Res. 213. https://doi.org/10.1016/j.still.2021.105143. 416 Amelung, W., Bossio, D., de Vries, W., Kögel-Knabner, I., Lehmann, J., Amundson, R., Bol, R., Collins, C., Lal, R., 417 Leifeld, J., Minasny, B., Pan, G., Paustian, K., Rumpel, C., Sanderman, J., van Groenigen, J.W., Mooney, 418 S., van Wesemael, B., Wander, M., Chabbi, A., 2020. Towards a global-scale soil climate mitigation strategy. Nat. Commun. 11:5427. https://doi.org/10.1038/s41467-020-18887-7. 419 420 Baldock, J.A., Oades, J.M., Nelson, P.N., Skene, T.M., Golchin, A., Clarke, P., 1997. Assessing the extent of 421 decomposition of natural organic materials using solid-state ¹³C NMR spectroscopy. Soil Res. 35(5): 422 1061-1084. https://doi.org/10.1071/S97004. 423 Bai, X.X., Tang, J., Wang, W., Ma, J.M., Shi, J., Ren, W., 2023. Organic amendment effects on cropland soil organic 424 carbon and its implications: A global synthesis. Catena. 231:107343. https://doi.org/ 425 10.1016/j.catena.2023.107343. 426 Benbi, D.K., Boparai, A.K., Brar, K., 2014. Decomposition of particulate organic matter is more sensitive to 427 temperature than the mineral associated organic matter. Soil Biol. Biochem. 70, 183-192. 428 https://doi.org/10.1016/j.soilbio.2013.12.032. 429 Cambardella, C.A., Elliott, E.J., 1992. Particulate soil organic-matter changes across a grassland cultivation sequen 430 ce. Soil Sci. Soc. Am. J. 56(3), 777-783. https://doi.org/10.2136/sssaj1992.0361599500560003001 431 <u>7x</u>. 432 Chenu, C., Angers, D.A., Barre, P., Derrien, D., Arrouays, D., Balesdent, J., 2019. Increasing organic stocks in 433 agricultural soils: Knowledge gaps and potential innovations. Soil Till. Res. 188, 41-52. 434 https://doi.org/10.1016/j.still.2018.04.011. 435 Cotrufo, M.F., Ranalli, M.G., Haddix, M.L., Six, J., Lugato, E., 2019. Soil carbon storage informed by particulate 436 mineral-associated 12(12). 989-996 organic matter. Nature Geosci. 437 https://doi.org/10.1038/s41561-019-0484-6. 438 Dennis, P.G., Miller, A.J., Hirsch, P.R., 2010. Are root exudates more important than other sources of rhizodeposits 439 in structuring rhizosphere bacterial communities? FEMS Microbiol Ecol, 72, 313-327. https:// 440 doi.org/10.1111/j.1574-6941.2010.00860.x 441 Eusterhues, K., Rumpel, C., Kögel-Knabner, I., 2005. Organo-mineral association in sandy acid forest soils: 442 importance of specific surface area, iron oxides and micropores. Eur. J Soil Sci. 56, 753-763. 443 https://doi.org/10.1111/j.1365-2389.2005.00710.x.





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Table 1
 Soil organic carbon (SOC), total nitrogen (TN), the ration of SOC to TN (SOC/TN), pH, crystalline iron oxides
 (Fe_d), non-crystalline iron oxides (Fe_o), aggregate size distribution and mean weight diameter (MWD) of water stable aggregates under no manure (Control), low manure (LM), high manure (HM) and high manure plus lime
 (HML) treatments.

T	SOC (g kg ⁻¹)		SOC/TN	pН	Fe_d	Fe _o (g kg ⁻¹) –	Aggregate size distribution (g g-1)			MWD
Treatments					(g kg ⁻¹)		>0.25 mm	0.05-0.25 mm	<0.05 mm	(mm)
Control	4.79c	0.67c	7.10b	4.80c	62.99a	1.81b	0.67c	0.28a	0.05a	0.98b
LM	7.91b	1.00b	7.94a	5.42c	59.98a	1.74b	0.71b	0.25a	0.04ab	1.09b
HM	10.37a	1.35a	7.69ab	6.11b	59.59a	2.25a	0.77a	0.19b	0.03b	1.22a
HML	10.86a	1.40a	7.76ab	7.08a	62.22a	2.29a	0.78a	0.18b	0.04ab	1.32a

Values are the means (n=3). Different lowercase letters after values in the same row indicate a significant difference among four manure treatments (P < 0.05).

591 **Table 2**

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592 Mass proportion, SOC concentration and contribution of the mineral-associated organic matter (MAOM) (<53 µm) 593 fraction, particulate organic matter (POM) (>53 µm) fraction under no manure (Control), low manure (LM), high 594 manure (HM) and high manure plus lime (HML) treatments.

_	Mass proportion (%)		SOC conc (g C kg		Contribution to total SOC/%		
Treatments	MAOM	POM	MAOM	POM	MAOM	POM	
Control	74.3a	16.6b	4.18c	0.44c	82.4a	8.8c	
LM	75.1a	17.2ab	5.61b	1.26b	71.0b	15.9b	
HM	73.9a	19.4a	6.81a	2.46a	65.8b	23.7a	
HML	72.5a	19.3a	7.09a	2.84a	65.4b	26.0a	

Values are the means (n=3). Different lowercase letters after values in the same row indicate a significant difference among four manure treatments (P < 0.05). Calculations follow the methodology described in Section 2.4.





Table 3 The contents of various C functional groups in CPMAS-13C-NMR spectra under no manure (Control), low manure (LM), high manure (HM) and high manure plus lime (HML) treatments. Alkyl (0-45 ppm); O-alkyl C (45-110 ppm) and high manure plus lime (HML) treatments. ppm); aromatic C (110-160 ppm) and carbonyl C (160-220 ppm).

Treatments	alkyl C (%)	O-alkyl C (%)	aromatic C	carbonyl C (%)	alkyl C/ O-alkyl C	aromaticity (%)	aromatic C/O- alkyl C	aliphatic C/ aromatic C	aliphaticity (%)
Control	25.4b	44.9c	15.3a	14.4a	0.57a	17.9a	0.34a	4.58b	82.1b
LM	26.2ab	47.4a	13.3b	13.1b	0.55b	15.3b	0.28b	5.53a	84.7a
HM	26.6a	46.1b	13.4b	13.9ab	0.58a	15.6b	0.29b	5.41a	84.4a
HML	26.5a	45.9b	13.3b	14.3a	0.58a	15.6b	0.29b	5.43a	84.4a

Values are the means (n=3). Different lowercase letters after values in the same row indicate a significant difference among four manure treatments (P < 0.05). Aromaticity = aromatic C/ (alkyl C + O-alkyl C + aromatic C); aliphatic C = (alkyl C + O-alkyl C); aliphaticity = (alkyl C + O-alkyl C) / (alkyl C + O-alkyl C + aromatic C).

Table 4

The mass losses during specific temperature ranges under no manure (Control), low manure (LM), high manure (HM) and high manure plus lime (HML) treatments. Exo1 represents thermally labile soil organic matter (SOM); Exo2 represents thermally stable SOM; Exotot represents total SOM. TG-T50 indicates the temperature at which half of the total SOM is lost.

Treatments	Exo ₁ (180~380°C) (%)	Exo ₂ (380~530°C) (%)	Exo _{tot} (180~530°C) (%)	Exo ₁ /Exo _{tot}	TG-T ₅₀ (°C)
Control	2.28c	2.36b	4.64b	0.49b	381a
LM	2.62b	2.54a	5.16a	0.51a	376ab
HM	2.69ab	2.68a	5.37a	0.50b	369b
HML	2.89a	2.56a	5.45a	0.53a	371b

Values are the means (n=3). Different lowercase letters after values in the same row indicate a significant difference among four manure treatments (P < 0.05).





Figure legends

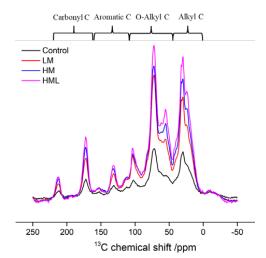
623 Fig. 1 CPMAS-13C-NMR spectra under no manure (Control), low manure (LM), high manure (HM) and high 624 manure plus lime (HML) treatments. Alkyl (0-45 ppm); O-alkyl C (45-110 ppm); aromatic C (110-160 ppm) and 625 carbonyl C (160-220 ppm) 626 627 Fig. 2 Thermogravimetry (TG) curves (a) and corresponding derivative thermogravimetry (DTG) curves (b) of soil 628 organic matter (SOM) under no manure (Control), low manure (LM), high manure (HM) and high manure plus 629 lime (HML) treatments 630 631 Fig. 3 The relationships between SOC and variables related to the chemical protection and physical protection 632 633 Fig. 4 Biplots of the principal component analysis (PCA) between the quantity and quality of SOC and variables 634 related to chemical protection, physical protection across four manure application treatments (CK, Control; LM, 635 low manure; HM, high manure; and HML, high manure plus lime) 636 637

Fig. 5 The relationships between SOC quantity, chemical recalcitrance and thermal stability

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 $\label{eq:Fig.1} \textbf{Fig.1} \ CPMAS-{}^{13}C-NMR \ spectra \ under \ no \ manure \ (Control), \ low \ manure \ (LM), \ high \ manure \ (HM) \ and \ high \ manure \ plus \ lime \ (HML) \ treatments. \ Alkyl \ (0-45 \ ppm); \ O-alkyl \ C \ (45-110 \ ppm); \ aromatic \ C \ (110-160 \ ppm) \ and \ carbonyl \ C \ (160-220 \ ppm)$

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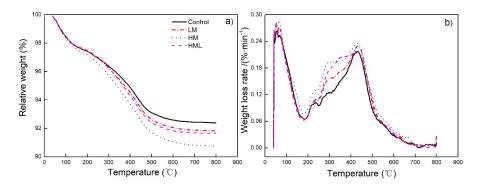


Fig. 2 Thermogravimetry (TG) curves (a) and corresponding derivative thermogravimetry (DTG) curves (b) of soil organic matter (SOM) under no manure (Control), low manure (LM), high manure (HM) and high manure plus lime (HML) treatments





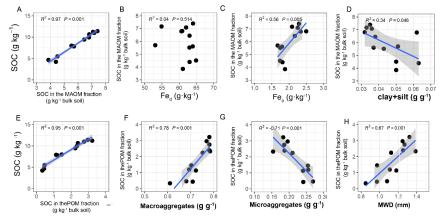


Fig. 3 The relationships between SOC and variables related to the chemical protection and physical protection. (A to D) Variables characterizing chemical protection, including the content of SOC stored in the MAOM fraction, the content of crystalline Fe oxides (Fe_0), the content of non-crystalline Fe oxides (Fe_0), and the content of clay and silt. (E to H) Variables characterizing physical protection, including the content of SOC stored in the POM fractions, the content of soil macroaggregates, the content of soil microaggregates and the mean weight diameter (MWD)





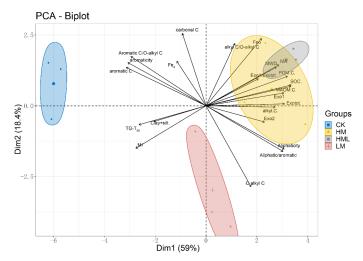


Fig. 4 Biplots of the principal component analysis (PCA) between the quantity and quality of SOC and variables related to chemical protection, physical protection across four manure application treatments (CK, Control; LM, low manure; HM, high manure; and HML, high manure plus lime). Fe_d, crystalline Fe oxides; Fe_o, non-crystalline Fe oxides; MA, macroaggregates (>0.25 mm); MI, microaggregates (0.05-0.25 mm); MWD, mean weight diameter; aromaticity, aromatic C / (alkyl C + O-alkyl C + aromatic C); aliphatic C, Alkyl C + O-alkyl C; aliphaticity, (Alkyl C + O-alkyl C) / (Alkyl C + O-alkyl C + Aromatic C); Exo₁, thermally labile soil organic matter (SOM); Exo₂, thermally stable SOM; Exo_{tot}, total SOM; TG-T₅₀, the temperature at which half of the total SOM was lost





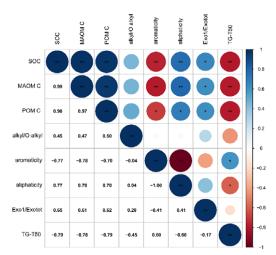


Fig. 5 The relationships between SOC quantity, chemical recalcitrance and thermal stability. Aromaticity, aromatic
C / (alkyl C + O-alkyl C + aromatic C); aliphaticity, (alkyl C + O-alkyl C) / (alkyl C + O-alkyl C + aromatic C);
Exo₁, thermally labile soil organic matter (SOM); Exo_{tot}, total SOM; TG-T₅₀, the temperature at which half of the
total SOM was lost