1 Long-term pig manure application increases SOC through aggregate protection and Fe-C

- 2 associations in a subtropical Red soil (Udic Ferralsols)
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Abstract

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Manure is known to improve soil organic carbon (SOC) in Fe-rich red soils, but the underlying stabilization mechanisms remain poorly understood. A long-term field experiment was conducted with four treatments: no amendment (Control), low manure (LM, 150 kg N ha⁻¹ yr⁻¹), high manure (HM, 600 kg N ha⁻¹ yr⁻¹), and high manure with lime (HML, 600 kg N ha⁻¹ yr⁻¹ plus 3000 kg Ca (OH)₂ha⁻¹ 3yr⁻¹). The quantity and quality of SOC were characterized by physical fractionation, 13 C NMR spectroscopy and thermogravimetry (TG) analysis from 0-20 cm depth. Manure application increased total SOC by 65.1%-126.7%, with the increase primarily occurring in the particulate organic matter (POM) fraction. Specifically, POM C increased substantially by 208%-592%, compared to a 32.0% -66.8% increase in MAOM C. POM C was stabilized through the process of hierarchical aggregation. Fresh manure inputs acted as binding nuclei, leading to an increase in macroaggregates (>0.25 mm) concurrently with a reduction in microaggregates (0.05–0.25 mm), physically isolating labile C from microbial decomposition. Concurrently, manure amendments triggered Fe-mediated chemical stabilization. The elevated pH (4.8 to 5.4-7.1) in the manure treatments enhanced non-crystalline Fe oxide (Fe₀) content by 25.4%, which positively correlated with MAOM C ($R^2 = 0.56$, P < 0.05). Despite chemical composition shift toward aliphaticity and

reduced aromaticity, thermally stable organic matters increased by 8%-12%, revealing the critical role of Feo in offsetting inherent molecular lability when aggregate protection was removed by destruction prior to TG analysis. Overall, this study proposes a dual SOC stabilization framework via physical and chemical Fe-carbon associations in a subtropical red soil.

Keywords: Particulate organic matter; Mineral-associated organic matter; Nuclear magnetic resonance; Thermogravimetry analysis

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1. Introduction

Soil organic carbon (SOC), the largest carbon reservoir in the terrestrial ecosystems, plays critical roles in climate mitigation and soil multifunctionality (Amelung et al., 2020; Lal, 2004). Red soils (Udic Ferralsols, according to Chinese Soil Taxonomy) have low SOC content due to intense weathering and rapid mineralization in tropical and subtropical South China (Yan et al., 2013; Zhang et al., 2013). While manure has been widely used to enhance SOC in these soils (Bai et al., 2023; Nichitha et al., 2023; Zhang et al., 2023), the underlying mechanisms governing its stabilization remain elusive. Currently, the discrepancies existed regarding the manure-induced SOC accrual through chemical recalcitrance, physical protection (by aggregation processes), or organo-mineral interactions—given the Fe-enrich red soils and their pH-dependent reactivity (Kleber et al., 2021; Six et al., 2002; Song et al., 2022). Considering, the red soils covering 22% of China's cropland and low inherent SOC levels, it is essential to unveil Fe-C interactions that regulate SOC stabilization through organic amendments. Existing studies showed some conflicting evidences on SOC stabilization pathways. For instance, Mustafa et al. (2021) reported increased aromatic C (chemically recalcitrant) with manure application, whereas Yan et al. (2013) observed preferential accumulation of labile O-alkyl C. This paradox highlighted uncertainties on how manure inputs altered SOC composition. It is reported that Fe oxides contributed to SOC stabilization in red soils (Zhang et al., 2013). These reactive Fe phases can form stable covalent bonds between their surface hydroxyls and organic functional groups, protecting SOC from microbial decomposition (Ruiz et al., 2024). In comparison with crystalline Fe oxides (Fe_d), non-crystalline Fe oxides (Fe_o) exhibit organic matter adsorption capacity primarily attributed to their larger specific surface area (Zhang et al., 2013). The ratio of

Fed and Fed is dynamically regulated by pH, which can be intentionally manipulated through manure application (Liu et al., 2020; Wang et al., 2023). The long term impacts of manure induced pH shifts on Fe oxide speciation and related organic carbon sequestration are still not well understood, despite the known connection between pH driven Fe oxide transformation and organic matter stabilization. Previous studies have demonstrated that long-term application of organic fertilizers—including manure, compost, and straw—significantly increased amorphous iron oxides (Fe₀) in agricultural soils (Chen et al., 2022; Huang et al., 2018; Wang et al., 2019). These studies further revealed that redox cycling promoted the reductive dissolution of crystalline iron (Fe_d), followed by its reprecipitation as amorphous Fe_o during oxidative periods. On the other hand, soil pH was identified as a key factor dynamically influencing the Fe_d/Fe₀ ratio (Liu et al., 2020; Wang et al., 2023). Although organic amendments are known to elevate pH in acidic red soils, the specific effect of soil pH modulating the formation and stabilization of Fe_o remains unclear. Additionally, previous studies widely isolated chemical recalcitrance, aggregation protection, or organo-mineral interactions separately, largely neglecting the integrative assessments including these pathways. To address this knowledge gap, an integrative approach combining physical fractionation, molecular characterization, and thermal stability analysis is essential for elucidating the coupled effects of manure on SOC quantity and quality. Physically separating soil organic matter (SOM) into mineral-associated organic matter (MAOM) and particulate organic matter (POM) fractions could help to predict SOC dynamics, and clarify SOC stabilization mechanisms. The POM associated SOC (defined as POM C) physically protected in the macroaggregates, while MAOM associated SOC (defined as MAOM-C) chemically protected via organo-mineral bonding (Chenu et al., 2019; Lavallee et al., 2019; Poeplau et al., 2018). Although physical fractionation could effectively isolate operationally defined these SOC pools (Poeplau et al., 2018), it failed to reveal the chemical heterogeneity of SOC (Cotrufo et al., 2019; Lavallee et al., 2019). Thus, solid-state ¹³C NMR spectroscopy could address this gap by quantifying SOC functional groups (e.g. alkyl, O-alkyl, aromatic C), after completely disrupting organo-mineral interaction by HF pretreatment (Kögel-Knabner, 1997). Additionally, the thermogravimetry (TG) could rapidly assess the thermal stability of SOC without requiring pretreatment (Gao et al., 2015). On the other hand, some advanced techniques such as nano-scale secondary ion mass spectrometry

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(NanoSIMS) and high-resolution mass spectrometry could identify the spatial visualization of organo-mineral associations and molecular characterization, these approaches heavily required complex instrumentation and are not cost-effective for the comparative analysis of multiple samples. Therefore, combined NMR with TG approaches to efficiently evaluate the chemical composition and thermal stability of SOC in response to manure management.

The specific objectives of this study were: 1) to evaluate the changes of Fe oxides and its effect on MAOM formation; 2) to explore how soil aggregation affected POM formation; 3) to evaluate the effect of manure application on SOC composition and stability. We hypothesized that: 1) Manure application could enhance MAOM C formation by increasing non-crystalline Fe oxides (Fe_o), induced by elevated pH; 2) Manure application would strengthen the physical protection due to the improved soil aggregation, which triggered labile SOC protection; 3) The application of pig manure would boost the recalcitrance of SOC, thus the thermal stability. The specific objectives of this studies were: 1) to evaluate the changes of Fe oxides and its effect on MAOM formation; 2) to explore how soil aggregation affected POM formation; 3) to evaluate the effect of manure application on SOC composition and stability.

2. Materials and methods

2.1. Site description and experimental design

The long-term field experiment is located at Yingtan National Agroecosystem Field Experiment Station of the Chinese Academy of Sciences (28°15′20″N, 116°55′30″E) in Jiangxi Province, China. The site has a typically subtropical humid monsoon climate with a mean annual temperature of 17.6°C and precipitation of 1795 mm (Jiang et al., 2018). The soil is derived from Quaternary red clay, and is classified as Udic Ferralsols according to Chinese Soil Taxonomy. The soil contains 36.3% clay, 45.2% silt and 21.2% sand.

The field experiment was initiated in April, 2002. The soils used in this experiment were transported from adjacent forestland and repacked (to a depth of 2 m) for all the plots. This design ensures a uniform initial soil condition and eliminates the influence of prior land use history. Four treatments were compared: (1) no manure amendment (Control), (2) low pig manure with 150 kg N ha⁻¹ a⁻¹ (LM), (3) high pig manure with 600 kg N ha⁻¹ a⁻¹ (HM), and (4) high pig manure with 600

kg N ha⁻¹ a⁻¹ and lime (HML). The four treatments received solely pig manure as the nitrogen source, with no synthetic fertilizers applied. The pig manure, collected from nearby pig farms, contained an average total carbon of 386.5 g kg⁻¹, total nitrogen of 36.2 g kg⁻¹ and total phosphorus of 21.6 g kg⁻¹ ¹ on a dry matter basis. The annual amount of pig manure applied to each treatment was calculated based on its nitrogen content. Since all aboveground residues (stalks and leaves) and manually recoverable roots were completely removed from the field after harvest, the total carbon inputs to the soil were derived mainly exclusively from pig manure. This resulted in average annual carbon inputs of around 1.6 and 6.4 Mg C ha-1 for the LM and HM treatments, respectively (see Supplementary Material for calculation details). The field experiment was set up following a completely randomized design, with each treatment has three replicate plots. Each plot has a size of $2 \text{ m} \times 2 \text{ m}$. Lime was applied at 3 000 kg Ca (OH)₂ ha⁻¹ (3a)⁻¹ for the HML treatment. The field was planted with corn (Zea mays L.) monoculture annually from April to July. The pig manure was applied annually each April onto the soil surface (0-10 cm depth) and subsequently incorporated into the topsoil through plowing and harrowing. All the management measures, including sowing, harvesting and weeding, were manually operated. 2.2 Sampling Sampling was conducted in July 2019, after the harvest of corn. Triplicate topsoil samples (0-20 cm) were randomly collected with a shovel from each plot and composited together to form one bulk sample. The soil samples were air-dried at room temperature and were gently crushed with a rubber mallet to pass through an 8-mm sieve, preserving aggregates >8mm for further analysis. Visible plant residues, roots and stones were removed (Soil Survey Staff, 2011). No visible manure residues were observed. 2.3 Soil properties measurements SOC and Total nitrogen (TN) were determined by an elemental analyzer (Vario MACRO, Elementar, Germany). Soil pH was measured using a glass electrode (PHS-3D, SANXIN, China) with soil: deionized water ratio of 1: 2.5. Crystalline (Fe_d) and non-crystalline Fe oxides (Fe_o) were extracted by DCB (Dithionite-citrate-bicarbonate) and oxalate, respectively (Yan et al., 2013), and then were determined by graphite furnace atomic absorption spectrometry (GFAAS) (PinAAcle

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900T, PerkinElmer, America). Water stability of aggregates was tested using the fast-wetting method

- 146 following Le Bissonnais (1996). Aggregate stability was expressed as mean weight diameter
- 147 (MWD). Detailed experimental processes and calculation can be found in Zhou et al. (2019).
- 148 2.4. Physical fractionation
- Soil was fractionated into MAOM (<53 μm) and POM (>53 μm) fraction following Cambardella
- and Elliott (1992) and Cotrofu et al. (2019). Briefly, 10 g sieved samples (<2 mm) were completely
- dispersed in dilute sodium hexametaphosphate ((NaPO₃)₆, 0.5%) at a soil: solution ratio of 1:4 by
- shaking for 18 h (25°C, 180 r min⁻¹). This procedure primarily disrupts macroaggregates (>250 μm),
- but remaining that stable microaggregates (<53 μm). After dispersing, soil slurry was passed through
- a 53 µm sieve and rinsed several times with deionized water. The fraction passing through the sieve
- was collected as MAOM fraction, and that remaining on the sieve was collected as POM fraction.
- Both fractions were centrifuged and the solution was decanted, and then the remaining material was
- oven-dried to constant weights at 60°C. SOC concentration of each fraction was measured using the
- wet oxidation method. Dried mass proportions of each fraction (g fraction g⁻¹ soil) were calculated
- as follows:

$$f_M = m_M / m_{bulk} \tag{1}$$

$$f_P = m_p / m_{bulk} \tag{2}$$

- where f_M and f_P were the dried mass proportions of MAOM and POM fraction (g fraction g^{-1} bulk
- soil), respectively; m_M and m_P (g)were the dried masses of MAOM and POM fractions; m_{bulk} (g)
- was the dried mass of bulk soil.
- SOC in the MAOM and POM fractions were called as MAOM C and POM C, respectively in
- this paper. MAOM C and POM C were calculated by multiplying the dried mass proportions of each
- fraction (g fraction g⁻¹ soil) by the respective SOC concentrations (g C kg⁻¹ fraction) as follows
- 168 (Garten and Wullschleger, 2000; Lian et al., 2015):

$$MAOM C = f_M \times SOC_M$$
 (3)

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$$POM C = f_P \times SOC_P \tag{4}$$

- where f_M and f_P were calculated by Equation (1) and Equation (2); SOC $_M$ and SOC $_P$ were the SOC
- concentrations in the MAOM and POM fraction (g C kg⁻¹ fraction), respectively.
- 173 The contributions of MAOM C and POM C to total SOC (%) was calculated as:
- 174 Contribution of MAOM C (%) = MAOM C / Total SOC \times 100 (5)

175 Contribution of POM C (%) = POM C / Total SOC \times 100 (6)

where MAOM C and POM C were derived from Equations (3) and (4), respectively; total SOC was
the content of SOC in the bulk soil.

2.5. SOC chemical composition and chemical stability

SOC chemical composition was analyzed with a solid-state cross-polarization magic angle spinning (CPMAS) ¹³C nuclear magnetic resonance (NMR) spectroscopy. Prior to NMR analysis, <2 mm air-dried soils were pretreated with hydrofluoric acid (HF) to remove paramagnetic Fe³⁺ (iron oxides) following Gao et al. (2021). While HF pretreatment can dissolve short-range-order minerals and potentially disrupt certain organo-mineral associations, it is an essential step to improve the signal-to-noise ratio and spectral resolution required for the reliable quantification of carbon functional groups. Firstly, 20 g soil was mixed with 100 mL 10% (w/w) HF solution in a polyethylene bottle and then shaken for 0.5 h per day for three days. Afterwards, the supernatant liquid was discarded and another 100 mL 10% HF solution was added again. The above procedures were repeated for 15 times. The residue was rinsed 10 times with deionized water until the pH was close to neutral. The remaining soil was freeze-dried and then ground in an agate mortar to pass through a 100-mesh sieve (0.149 mm) for further analysis. This fine grinding ensured homogeneous packing in the NMR rotor, minimizing signal heterogeneity (Simpson & Simpson, 2012).

Carbon functional groups were determined with the Bruker Ascend 500 MHz NMR spectrometer (Bruker BioSpin, Rheinstetten, Germany). Dry powdered samples were placed in a 4-mm sample rotor operating at a ¹³C resonance frequency of 125.8 MHz. The NMR spectrometer run at a spinning rate of 5kHz, and 10500 scans were collected for each sample. The spectra were collected over an acquisition time of 12 ms and a recycle delay of 0.8 s. All obtained spectra were processed with phase and baseline corrections. To allow for a direct comparison of the relative distribution of carbon functional groups, the total spectral intensity of each spectrum was normalized to that of the control soil sample. The normalized spectra were then integrated across and we assigned the obtained spectra to four different carbon functional groups, i.e., alkyl C (0-45 ppm), O-alkyl C (45-110 ppm), aromatic C (110-160 ppm) and carbonyl C (160-220 ppm) according to Kögel-Kanbner (1997). The relative concentrations of the different functional groups were calculated as the percentage of their peak areas to the total aeras using MestReNova 14.0 software (Mestrelab Research, 2019, Spain).

Indices used to evaluate SOC recalcitrance include: alkyl C/O-alkyl C, aromaticity, aromatic C/O-alkyl C, aliphatic C/aromatic C and aliphaticity (Baldock et al., 1997; Du et al., 2017). Aromaticity = aromatic C / (alkyl C + O-alkyl C + aromatic C); Aliphatic C = alkyl C + O-alkyl C; Aliphaticity = (alkyl C + O-alkyl C) / (alkyl C + O-alkyl C + aromatic C). The alkyl C/O-alkyl C ratio reflects the degree of microbial transformation, with higher values indicating advanced decomposition and accumulation of recalcitrant alkyl compounds (Baldock et al., 1997). The aromatic C/O-alkyl C ration aligns with the alkyl C/O-alkyl C ration in reflecting the degree of SOC decomposition, whereas the aliphatic C/aromatic C ratio presents a contrary perspective to them. Aromaticity is a chemical concept denoting the resistance to microbial degradation (Kögel-Knabner, 1997). Aliphaticity is used to quantify the proportion of labile aliphatic components relative to stable aromatic moieties.

2.6. Thermogravimetry analysis

TG analysis was performed using a Netzsch TG 209F1 (Netzsch-Gerätebau GmbH, Selb, Germany). Air-dried soil samples were first sieved through a 2-mm mesh, and then ground in a ball mill to pass through a 50 μ m sieve. The grounded soil samples (5 ~ 10 mg) were placed in an Al₂O₃ crucible covered with an aluminum lid and were oxidized in an atmosphere of 20 mL min⁻¹ of synthetic air (20% O₂ and 80% N₂) and 20 mL min⁻¹ of N₂ as a protective gas. The temperature program included a heating rate of 10 $^{\circ}$ C min⁻¹ from 40 $^{\circ}$ C up to 800 $^{\circ}$ C. The sample mass percentage relative to the initial mass as a function of temperature was recorded simultaneously, and its first derivative (DTG) was calculated to represent the mass loss rate.

Three processes were detected in the temperature range of 40°C to 800°C: hygroscopic moisture evaporation, SOM decomposition and carbonate breaking down (Gao et al., 2015; Siewert, 2004). Based on the observed DTG curve, the weight loss between 180°C and 530°C, representing the temperature range during which SOM was decomposed, was defined as the total mass of SOM (Exotot). The mass loss between 180°C and 380°C was a consequence of thermally labile SOM oxidation (Exo1) and that between 380°C and 530°C was caused by the combustion of more stable organic matter (Exo2) (Gao et al., 2015; Volkov et al., 2020). We used two parameters, the ratio of Exo1 and Exotot (Exo1/Exotot) and the temperature at which half of the SOM was decomposed (TG-T50), to characterize the thermal stability of SOM. Higher Exo1/Exotot and lower TG-T50 values

indicate that the sample has more thermally labile or unstable SOM (Gao et al., 2015; Siewert, 2004).

2.7. Statistical analysis

Statistical analyses were carried out with R Studio software (R Development Core Team, version 4.1.2). One-way analysis of variance (ANOVA) was conducted to assess the effect of amendments on soil physico-chemical properties, SOC physical fractions, chemical composition and thermal indices. The independence of samples, normality of residuals residues and homogeneity of variances were checked by Chisq, Shapiro-Wilk and Bartlett test, respectively. Fisher's least significant difference (LSD) method was used for the multiple comparisons of means with a 0.05 significance level. Linear regression analysis was conducted to investigate the correlations between SOC and iron oxides, soil aggregation. Principal component analysis (PCA) was performed to evaluate the relationship between the quantity and quality of SOC and factors related to chemical protection and physical protection. Pearson correlation analyses were performed to explore the relationships between chemical composition and thermal indices.

3. Results

247 3.1. Long-term pig manure application increased SOC, TN, pH, non-crystalline (Fe_o) and improved soil aggregation

Long-term manure amendment altered the soil physico-chemical physic chemical properties (Table 1). Relative to the Control, the SOC concentration was LM, HM and HML treatments increased SOC concentration by 64.9%, 116.2% and 126.6% higher in the LM, HM and HML treatments, respectively (P < 0.05), and increased TN concentration by was 48.0%-108.2% higher (P < 0.05). The pH values were increased by 0.62-2.28 units higher than those in the Controlafter pig manure application (P < 0.05). Application of pig manure had no significant effect on crystalline iron oxides (Fe_d) content (P > 0.05), but the HM and HML treatments showed significantly greater increased non-crystalline iron oxides (Fe_o), with an increase of by 25.4% relative to the Control (P < 0.05). Relative Compared to the Control, the HM and HML treatments had significantly higher increased macroaggregates (>0.25 mm) content by 15.8% and 16.8%, respectively, and they increased greater mean weight diameter (MWD) by 24.3% and 35.0%, respectively (P < 0.05). In contrast, Mmicroaggregates (0.05-0.25 mm) was decreased by 30.4% and 36.4%, respectively (P < 0.05).

263 The distribution of MAOM and POM fractions was significantly influenced by manure and lime 264 amendments (P < 0.05, Table 2). Across all treatments, MAOM dominated the soil mass proportion 265 (72.5%–75.1%), whereas the proportion of POM was mass proportion increased progressively higher, ranging from 16.6% in the eControl to 19.3%–19.4% under the HM and HML treatments. 266 267 Manure application improved SOC concentration in both fractions, SOC in the MAOM fraction 268 was-with increase rates of 32.0%-66.8% greater, and that in the POM fraction was 208%-592% 269 greater relative to the Controlin the MAOM and POM fractions, respectively. Despite this, 270 The contribution of MAOM to total SOC was lowerdeclined from 82.4% in the HM and HML 271 treatments (65.5%-65.8%) than in the Ceontrol (82.4%) to 65.5% 65.8% under the HM and HML treatments, while the contribution of POM was contribution increased nearly threefold higher 272 273 (increasing from 8.8% to 23.7%–26.0%). Lime addition (HML vs. HM) did not significantly alter 274 mass proportions but further enhanced POM C concentration (+15.4%) and its contribution to total 275 SOC-contribution (+9.7%). 276 3.3. Effect of long-term pig manure application on SOC chemical composition and recalcitrance The solid-state ¹³C NMR spectra showed different signal patterns for the different treatments (Fig. 277 278 1) and quantified the ratios of the different SOC functional groups shown in Table 3. Relative to the 279 Control, the HM and HML treatments had a significantly higher increased proportion of alkyl C by 280 4.6%-4.9% (P < 0.05), while the LM treatment showed no significant differencechange (P > 0.05). 281 The proportion of O-alkyl C was also significantly higher under the LM, HM and HML treatments, with increases of increased by 5.5%, 2.6% and 2.2% under the LM, HM and HML treatments, 282 283 respectively (P < 0.05). In contrast, the proportions of Aaromatic C and carbonyl C were lower in 284 the manure-amended treatments, with reduction of decreased by 12.4%-13.2% and 0.9%-8.7%, 285 respectively, in the manured treatment (P < 0.05). While aromatic C was increased in the content 286 after manure amend, its relative proportion was decreased by 12.4% - 13.2% (P < 0.05). 287 Relative to the Control, the LM treatment significantly decreased alkyl C/O-alkyl C ratio was 288 significantly lower under the LM treatment (P < 0.05), but whereas the HM and HML treatments 289 had no significant effect on the ratio (P > 0.05, Table 3). The aromaticity was decreased by 14.4%, 290 12.9% and 13.2% under LM, HM and HML treatments, respectively (*P* < 0.05, Table 3). Similarly,

3.2. Long-term pig manure application affected SOM physical fractions: MAOM and POM

291 the aromatic C/O-alkyl C ratio was decreased by 14.7%-17.7% lower in the manured treatment (P 292 < 0.05, Table 3). Conversely In contrast, both the aliphatic C/aromatic C ratio and aliphaticity were 293 increased by 18.1% 20.7% and 2.44% 3.66% higher in the manured treatments, with elevations of 18.1% - 20.7% and 2.44% - 3.66%, respectively (P < 0.05, Table 3). 294 295 3.4. The effect of long-term pig manure application on SOC thermal stability 296 The shapes of TG and its first derivatives (DTG) curves showed distinct weight losses rate as 297 temperature increased to above 100°C, in the order of HM>HML>LM>Control (Fig. 2). Relative to 298 the Control treatment, pig manure application significantly increased the total mass losses in the 299 range of 180°C to 530°C (Exo_{tot}) by 11.1%- 17.4% (P < 0.05, Table 4), and significantly increased 300 the mass losses during 180-380°C (Exo₁) and 380-530°C (Exo₂) by 14.6%-26.5% and 7.8%-13.7%, 301 respectively (P < 0.05, Table 4). The LM and HML treatments significantly increased the ratio of 302 $\text{Exo}_1/\text{Exo}_{\text{tot}}$ was significantly elevated in the LM and HM treatments (P < 0.05), whereas the HM 303 treatment showed had no significant effect relative to the Control (P > 0.05, Table 4). HM and HML 304 treatments significantly decreased TG-T₅₀ by 10.7-12.0 °C (P < 0.05), while LM treatments showed 305 no significant difference compared to Control (P > 0.05, Table 4). 306 3.5. Relationships between SOC and factors related to the mineral protection and physical 307 protection of SOC 308 Fig. 3 showed correlations between SOC and possible variables associated with the chemical 309 protection and physical protection of SOC. SOC in the bulk soil was significantly positively 310 correlated with MAOM C (R^2 =0.97, Fig. 3A). MAOM C showed no relationship with Fe_d (R^2 =0.04, 311 Fig. 3B), it was positively correlated with Fe_o (R^2 =0.56, Fig. 3C). A weaker correlation was observed 312 between MAOM C and the content of clay and silt $(R^2 = 0.34, P = 0.46, Fig. 3D)$ there was no 313 correlation between MAOM C and the content of clay and silt (R²=0.34, Fig.3D). Notably, Fe₀ 314 concentration was significantly and positively correlated with pH (R^2 =0.59, P<0.01; Fig. 4A). 315 Conversely, the concentration of Fe_d was not correlated with pH (P>0.05; Fig. 4B). 316 SOC in the bulk soil was significantly positively correlated with POM C (R^2 =0.95, Fig. 3E). POM 317 C was significantly associated with soil aggregation, evidenced by the positive correlation with 318 macroaggregates (>0.25 mm) (R^2 =0.78, Fig. 3F) and MWD (R^2 =0.67, Fig. 3H), while negative correlation with microaggregates (<0.25 mm) ($R^2=-0.71$, Fig. 3G). 319

A PCA plot diagram (Fig. 54) revealed distinct associations between SOC quantity/quality and stabilization mechanisms. SOC vector aligned positively with POM C, macroaggregates and MWD, but inversely with microaggregates. Chemically, SOC covaried with MAOM C and Fe_o, yet formed an obtuse angle (>90°) with Fe_d. TG-T₅₀ negatively associated with the proportion of O-alkyl C and aliphaticity (angles > 90°), but positively linked to aromatic C and aromaticity (angles < 90°).

4. Discussion

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4.1. POM C and physical protection

Long-term pig manure application caused an increase of SOC in the bulk soil relative to the Control, and the increase was mainly derived from continuous manure inputs (Gong et al., 2009). Although maize rhizodeposition (including root exudates and sloughed-off cells) (typically <10% of net primary productivity; Pausch and Kuzyakov, 2018) contributes to SOC during the growing season, its annual carbon inputs are only 0.2-0.5 Mg C ha⁻¹ yr⁻¹ (Dennis et al., 2010). This plantderived C input is much lower than that from manure, contribution is negligible in comparison with manure inputs, which ranged from 1.6 to 6.4 Mg C ha⁻¹ yr⁻¹ in the LM and HM treatments, respectively. The rigorous removal of all harvest residues (stalks and recoverable roots) minimizedexeluded aboveground and root biomass as a source of soil carbon deposition. It should be noted that some residual plant-derived materials (e.g., fine root fragments and exudates) have contributed marginally to the SOC pool. Given the relatively small quantity of plant-derived carbon inputs, the changes in SOC observed in this study are primarily attributable to the application of pig manure. Thus, the SOC increase in the LM/HM treatments (Table 1) can be predominately attributed to the exogenous organic C supplied by pig manure, and the HML treatment can be partly influenced by lime. Following microbial decomposition and transformation, the applied manure-C was progressively partitioned into distinct SOC pools. Manure-derived carbon was preferentially allocated to the POM fraction rather than the MAOM fraction. The contribution of POM C to SOC in the manure-amended treatment increased was nearly threefold as that in the Control (8.8% to 26.0%), while that the proportion of MAOM C decreased significantly from 82.4% in the Control to 65.4%–71.0% in the under manure application treatments (Table 2). The result was in line with Lan et al. (2022) and Li et al. (2018), who reported that increased manure substitution improved the POM C/MAOM C ratio, indicating manure was

beneficial to the formation of POM C. These results implied carbon inputs are preferentially stabilized as POM rather than MAOM. This finding is supported by biomarker evidence showing a great contribution of plant-derived carbon to the POM fraction (Zou et al., 2023). Consistent with this, a ¹³C isotopic tracing study revealed that 70%-87% of residue-derived SOC was accumulated in the POM fraction at two experimental sites (Mitchell et al., 2021). The mechanism behind this preferential sequestration is the physical occlusion of fresh organic material within macroaggregates, which provides initial protection against rapid microbial decomposition. Therefore, POM C can be a good indicator of SOC dynamics under field management (Álvaro-Fuentes et al., 2021; Wu et al., 2023). Recent biomarker evidence demonstrated that plant derived carbon contributes disproportionately to POM in comparison with MAOM fraction (Zou et al., 2023). Notably, 43C isotopic tracing study revealed that 70%-87% of residue-derived SOC was accumulated in the POM fraction at two sites in the study of Mitchell et al. (2021). These results jointly validate that fresh organic inputs are distributed primarily in the POM fraction due to direct occlusion within macroaggregates before microbial processing. Therefore, POM C can be a good indicator of SOC dynamics under field management (Álvaro-Fuentes et al., 2021; Wu et al., 2023). While POM C is inherently labile and susceptible to decomposition, it can be physically protected from rapid decay through occlusion within stable soil aggregates. POM C was susceptible to decomposition due to its rapid dynamics, whereas it was simultaneously protected through physical occultation within soil aggregates. According to the classical hierarchical aggregation model (Six et al., 2000), organic carbon acts as a primary binding agent for microaggregates (<0.25mm) to form macroaggregates (>0.25mm), with POM serving as both a structural nucleus and a transient carbon reservoir (Six et al., 2000). In this study, macroaggregates were significantly increased, and free microaggregates were significantly decreased after pig manure application (Table 1). These results implied that manure-derived carbon acted as binding agent, promoting the formation of macroaggregates from free microaggregates, bound microaggregates to form macroaggregates, thus This process provides physical protection for SOC, particularly that in the POM fraction, within the newly formed aggregates, providing physical protection for SOC particularly in the POM fraction (Peng et al., 2023). The observed positive correlation between POM C and both macroaggregates and aggregate MWD (Fig.3; Fig.4) further suggested verified the critical role of soil aggregation.

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However, the physical mechanisms by which POM facilitates this process (e.g., microbial mediation, hydrophobic interactions, or polysaccharide bridging) require further investigation. Experimental approaches such as isotopic labelling combined with micro-scale imaging (e.g., electron microscopy or X-ray computed tomography) can visualize the spatial distribution of POM within aggregates and quantify its role in aggregate formation. Future study should pay attention to these new technologies.

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4.2. MAOM C and chemical protection

MAOM retained higher SOC concentrations (4.18-7.09 g C kg⁻¹ bulk soil) than the POM fraction did, irrespective of treatments, though its contribution to total SOC decreased after manure application (Table 2). MAOM C remained the dominant SOC reservoir, consistently comprising over 65% of the total SOC across all treatments (Table 2). This underscores that organo-mineral complexation, rather than physical protection within aggregates, is the primary stabilization pathway in these red soils, even under significant organic input. MAOM C exhibited significantly longer mean residence time (MRT) compared to POM C, with reported turnover periods of 26-40 years versus 2.4-4.3 years, respectively (Benbi et al., 2014; Garten and Wullschleger, 2000). This fundamental difference originates from the unique stabilization mechanism of MAOM C through persistent organo-mineral associations (Lavallee et al., 2019). While clay minerals are widely recognized as key MAOM stabilizers (Hemingway et al., 2019; Liang et al., 2017), our study reveals a distinct iron oxide-dominated mechanism in these Fe-rich red soils. The strong correlation between MAOM C and non-crystalline Fe oxides (Fe_o, R^2 =0.56; Figs. 3C, 4) highlighted the pivotal role of Fe_o on SOC stabilization in this red soil. This iron-mediated stabilization likely stems from the exceptionally high specific surface area of Fe_o (~800 m² g⁻¹ in ferrihydrite; Kleber et al., 2005) and its superior capacity to form stable organo-mineral complexes (Eusterhues et al., 2005; Lehmann and Kleber, 2015). Previous studies have demonstrated that organic amendments can enhance the content of non-crystalline Fe oxides, thereby promoting Fe-C associations (Chen et al., 2022; Huang et al., 2017; Wang et al., 2019). Extending these findings, our results provide new mechanistic insight by demonstrating that the increase in non-crystalline Fe oxides is closely linked to manureinduced soil alkalization. Specifically, the significant positive correlation between pH and Feo

indicates that the The-increased pH after manure application (Table 1) created conditions favoring Fe_o preservation (Vithana et al., 2015), with an increase of 25.4% in the content of Fe_o. The higherinereased Fe_o content provided abundant reactive surfaces for MAOM formation. The nature of this interaction likely involves specific molecular-scale binding mechanisms. Recent advances in synchrotron-based spectroscopy (e.g., STXM-NEXAFS) have provided direct evidence that poorly crystalline iron oxides (e.g., ferrihydrite) stabilize organic carbon primarily through ligand exchange, where carboxyl (-COOH) and hydroxyl (-C-OH) functional groups in organic molecules replace the surface hydroxyl groups on Fe oxides (Ruiz et al., 2024). Furthermore, coprecipitation, where organic matter is encapsulated within the matrix of forming Fe (oxyhydr)oxides, represents another potent stabilization mechanism that can generate long-term sinks for organic carbon (Chen et al., 2014). While our bulk data cannot unequivocally distinguish between these mechanisms, the correlation suggests that similar processes are likely operative in our system, contributing to the stability of the large MAOM-C pool. This pH-Fe_o-MAOM nexus establishes a self-reinforcing stabilization mechanism: manure-derived organic ligands interact with Fe_o to form stable complexes, simultaneously protecting both organic carbon and Fe_o from dissolution (Kleber et al., 2021).

The positive correlation between Fe_o and MAOM C (Fig. 3C) suggested that non-crystalline Fe oxides played a critical role in stabilizing SOC through organo-mineral interactions. However, while our data support the association between Fe_o and MAOM C, the underlying mechanisms (e.g., adsorption, co-precipitation, or ligand exchange) remain speculative due to the lack of direct molecular-scale evidence. Future studies employing advanced spectroscopic techniques (e.g., synchrotron-based X-ray absorption spectroscopy or NanoSIMS) could explicitly characterize the binding forms of Fe-organic complexes, thereby validating the causal relationship between Feo and MAOM formation.

4.3. Chemical composition and thermal stability

Input of new organic material could alter chemical composition of SOC and lead to the change of the molecule recalcitrance of SOC (Guo et al., 2019; Yan et al., 2013; Zhou et al., 2010; Zhang et al., 2013). In the current study, pig manure application resulted in a higher proportion increased the content of O-alkyl C but a lower proportion of aromatic C, but declined that of aromatic C. Given Ppig manure was rich in cellulose and lignin components, so the its introduction of manure

436 greatly substantially increased O-alkyl C relative to the Control (Li et al., 2015). The considerable 437 accumulationinerease of O-alkyl C accounted for the reductiondeerease of the relative proportion 438 of aromatic C and the decrease of lower aromaticity. 439 Long-term pig manure application strengthened SOC thermal stability by improving the content 440 of thermally stable organic matters while it decreased TG-T₅₀. The observed increase in the absolute 441 amount of thermally stable SOC (Exo₂) can be attributed to the formation of stable organo-mineral 442 complexes with newly formed Fe_o. As elucidated by Kleber et al. (2021), the thermal stability of 443 organic matter is significantly enhanced upon binding to mineral surfaces. Thermal analysis 444 suggested that manure amend significantly increased SOM content, evidenced by the increase of 445 the mass losses during 180-530 °C (Exoto) (Table 4). The result was consistent with the change of 446 SOC content measured by conventional method (Table 1), indicating TG technology was promising 447 for measuring SOM content (Siewert, 2004; Tokarski et al., 2018). The decrease of TG-T₅₀ reflected 448 more easily decomposable SOC accumulated after manure application (Gao et al., 2015; Siewert, 449 2004). The result consisted with that in the NMR spectroscopy, where greater O-alkyl C and higher 450 aliphaticity were found under manure treatments (Table 3). SOC with more O-alkyl C functional 451 groups or higher aliphaticity was less likely to resist thermochemical degradation as revealed by the 452 negative relationship between TG-T₅₀ and aliphacity (Fig. 4; Fig. 5) (Hou et al., 2019; Lehman and Kleber, 2015), thus leading to a decrease of TG-T₅₀. In this study, the thermally labile organic 453 454 matters accounted for over half of the total organic matters, so the decrease of TG-T₅₀ after manure 455 application was just a result of the increased thermally organic matters not the decrease of thermal 456 stability. In contrast, thermal stability should be strengthened according to the increased thermally 457 stable organic matters after manure application. The observed decrease TG-T₅₀ and the simultaneous 458 increase in Exo₂ are not contradictory. This apparent paradox can be explained by the dual effect of 459 manure amendment: the input of a large quantity of labile, manure-derived organic compounds 460 significantly increased the proportion of thermally labile C, thereby reducing the average stability 461 of the entire SOM pool (as reflected by the decreased TG-T₅₀). Concurrently, the manure-induced 462 biogeochemical changes (e.g., increased pH and Fe oxide transformation) promoted the formation 463 of highly stable organo-mineral complexes. While the relative proportion of this stable C might be diluted by the new labile C, its absolute amount within the soil system increased substantially, which 464

is captured by the increase in the absolute Exo₂ signal. Since soil structure was destroyed due to the ground process of soil samples before TG analysis, the increase of the thermally stable organic matters was ascribed to mineral protection, where the correlations between organic carbon and Fe oxides increased the thermally resistance of OC (Ruiz et al., 2023).

5. Conclusion

Manure application increased SOC quantity and improved its quality in Fe-rich red soils. SOC in the POM fraction exhibited the most pronounced response to manure inputs, while the majority of SOC was stored in the MAOM fraction. Furthermore, SOC was stabilized by distinct yet complementary mechanisms: physical protection via aggregation process, and chemical protection via Fe-organic associations induced by elevated pH. In addition, manure application increased thermally stable organic matters. However, to better understand inherent mechanisms, future work should focus on molecular-scale characterization of Fe-organic interactions using synchrotron techniques (e.g., Fe K-edge XANES/EXAFS), and in-situ visualization of POM-mediated aggregation through advanced imaging tools (e.g., SEM-TEM or μ -CT) coupled with 13 C-labelled manure to track POM dynamics within aggregates.

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Author contributions

HR conceived the experimental approach, took soil samples from the field, carried out the laboratory and data analyses, wrote the first draft of the manuscript, and contributed to subsequent drafts. ZLD helped analyze NMR spectrum and revise the manuscript. WDG contributed to interpretating TG data and revising the manuscript. LXM conducted the experiment of NMR. XHP helped design the experiment, and analyze data. YJJ contributed to the design of field experiment and the process of taking soil samples. DMY contributed to the process of writing. HZ contributed to funding acquisition, conceiving the experimental approach, carrying out data interpretations, and the writing

494 of all subsequent manuscript drafts. 495 496 **Conflict of Interest** 497 Hu Zhou is a member of the editorial board of SOIL. 498 499 References 500 Álvaro-Fuentes, J., Franco-Luesma, S., Lafuente, V., Sen, P., Usón, A., Cantero-Martínez, C., and Arrúe, J.L.: Stover 501 management modifies soil organic carbon dynamics in the short-term under semiarid continuous maize, 502 Soil Till. Res., 213, 105143, https://doi.org/10.1016/j.still.2021.105143, 2021. 503 Amelung, W., Bossio, D., de Vries, W., Kögel-Knabner, I., Lehmann, J., Amundson, R., Bol, R., Collins, C., Lal, R., 504 Leifeld, J., Minasny, B., Pan, G., Paustian, K., Rumpel, C., Sanderman, J., van Groenigen, J.W., Mooney, 505 S., van Wesemael, B., Wander, M., and Chabbi, A.: Towards a global-scale soil climate mitigation strategy. 506 Nat. Commun., 11, 5427, https://doi.org/10.1038/s41467-020-18887-7, 2020. Baldock, J.A., Oades, J.M., Nelson, P.N., Skene, T.M., Golchin, A., and Clarke, P.: Assessing the extent of 507 508 decomposition of natural organic materials using solid-state ¹³C NMR spectroscopy, Soil Res., 35, 1061– 509 1084, https://doi.org/10.1071/S97004, 1997. 510 Bai, X.X., Tang, J., Wang, W., Ma, J.M., Shi, J., and Ren, W.: Organic amendment effects on cropland soil organic carbon and its implications: A global synthesis, Catena., 231, 107343, https://doi.org/ 511 512 10.1016/j.catena.2023.107343, 2023. 513 Benbi, D.K., Boparai, A.K., and Brar, K.: Decomposition of particulate organic matter is more sensitive to 514 temperature than the mineral associated organic matter, Soil Biol. Biochem., 70, 183-192, 515 https://doi.org/10.1016/j.soilbio.2013.12.032, 2014. 516 Cambardella, C.A., and Elliott, E.J.: Particulate soil organic-matter changes across a grassland cultivation sequenc 517 e, Soil Sci. Soc. Am. J., 5, 777-783, https://doi.org/10.2136/sssaj1992.03615995005600030017x, 518 1992. 519 Chen, M., Zhang, S., and Liu, L.: Organic fertilization increased soil organic carbon stability and sequestration by i 520 mproving aggregate stability and iron oxide transformation in saline-alkaline soil. Plant Soil 474, 233-24 521 9. https://doi.org/10.1007/s11104-022-05326-3, 2022. 522 Chenu, C., Angers, D.A., Barre, P., Derrien, D., Arrouays, D., and Balesdent, J.: Increasing organic stocks in 523 agricultural soils: Knowledge gaps and potential innovations, Soil Till. Res., 188, 41-52, 524 https://doi.org/10.1016/j.still.2018.04.011, 2019. Cotrufo, M.F., Ranalli, M.G., Haddix, M.L., Six, J., and Lugato, E.: Soil carbon storage informed by particulate and 525 526 mineral-associated organic matter, Nature Geosci., 12, 989-996, https://doi.org/10.1038/s41561-019-527 0484-6, 2019. 528 Dennis, P.G., Miller, A.J., and Hirsch, P.R.: Are root exudates more important than other sources of rhizodeposits in 529 structuring rhizosphere bacterial communities? FEMS Microbiol Ecol., 72, 313-327, https:// 530

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Table 1578 Soil organic carbon (SOC), total nitrogen (TN), the ration of SOC to TN (SOC/TN), pH, crystalline iron oxides (Fe₀), non-crystalline iron oxides (Fe₀), aggregate size distribution and mean weight diameter (MWD) of water-stable aggregates under no manure (Control), low manure (LM), high manure (HM) and high manure plus lime (HML) treatments.

| Tuestus sute | SOC (g kg ⁻¹) | TN (g kg ⁻¹) | SOC/TN | pН | Fe_d | Fe _o (g kg ⁻¹) | Aggregate size distribution (g g ⁻¹) | | | MWD |
|--------------|------------------------------|-----------------------------|--------|-------|-----------------------|---------------------------------------|--|-----------------|----------|-------|
| Treatments | | | | | (g kg ⁻¹) | | >0.25 mm | 0.05-0.25 mm | <0.05 mm | (mm) |
| Control | 4.79c | 0.67c | 7.10b | 4.80c | 62.99a | 1.81b | 0.67c | 0.28a | 0.05a | 0.98b |
| LM | 7.91b | 1.00b | 7.94a | 5.42c | 59.98a | 1.74b | 0.71b | 0.25a | 0.04ab | 1.09b |
| HM | 10.37a | 1.35a | 7.69ab | 6.11b | 59.59a | 2.25a | 0.77a | 0.19b | 0.03b | 1.22a |
| HML | 10.86a | 1.40a | 7.76ab | 7.08a | 62.22a | 2.29a | 0.78a | 0.18b | 0.04ab | 1.32a |

Values are the means (n=3). Different lowercase letters after values in the same row indicate a significant difference among four manure treatments (P < 0.05). Fed: crystalline oxides; Feo: non crystalline oxides.

Table 2 Mass proportion, SOC concentration and contribution of the mineral-associated organic matter (MAOM) ($<53 \mu m$) fraction, particulate organic matter (POM) ($>53 \mu m$) fraction under no manure (Control), low manure (LM), high manure (HM) and high manure plus lime (HML) treatments.

| | Mass proportion (%) | | SOC conc (g C kg | | Contribution to | Contribution to total SOC/% | |
|------------|---------------------|--------|---------------------|-------|-----------------|-----------------------------|--|
| Treatments | MAOM | POM | MAOM | POM | MAOM | POM | |
| Control | 74.3a | 16.6b | 4.18c | 0.44c | 82.4a | 8.8c | |
| LM | 75.1a | 17.2ab | 5.61b | 1.26b | 71.0b | 15.9b | |
| HM | 73.9a | 19.4a | 6.81a | 2.46a | 65.8b | 23.7a | |
| HML | 72.5a | 19.3a | 7.09a | 2.84a | 65.4b | 26.0a | |

Values are the means (n=3). Different lowercase letters after values in the same row indicate a significant difference among four manure treatments (P < 0.05). Calculations follow the methodology described in Section 2.4.

Table 3
 The contents of various C functional groups in CPMAS-¹³C-NMR spectra under no manure (Control), low manure
 (LM), high manure (HM) and high manure plus lime (HML) treatments. Alkyl (0-45 ppm); O-alkyl C (45-110 ppm); aromatic C (110-160 ppm) and carbonyl C (160-220 ppm).

| Treatments | alkyl C (%) | O-alkyl C (%) | aromatic C (%) | carbonyl C (%) | alkyl C/ O-alkyl C | aromaticity (%) | aromatic C/O- alkyl C | aliphatic C/ aromatic C | aliphaticity (%) |
|------------|----------------|------------------|----------------|----------------------|--------------------------|-----------------|-----------------------------|----------------------------|---------------------|
| Control | 25.4b | 44.9c | 15.3a | 14.4a | 0.57a | 17.9a | 0.34a | 4.58b | 82.1b |
| LM | 26.2ab | 47.4a | 13.3b | 13.1b | 0.55b | 15.3b | 0.28b | 5.53a | 84.7a |
| HM | 26.6a | 46.1b | 13.4b | 13.9ab | 0.58a | 15.6b | 0.29b | 5.41a | 84.4a |
| HML | 26.5a | 45.9b | 13.3b | 14.3a | 0.58a | 15.6b | 0.29b | 5.43a | 84.4a |

Values are the means (n=3). Different lowercase letters after values in the same row indicate a significant difference among four manure treatments (*P* < 0.05). Aromaticity = aromatic C/ (alkyl C + O-alkyl C + aromatic C); aliphatic C = (alkyl C + O-alkyl C); aliphaticity = (alkyl C + O-alkyl C) / (alkyl C + O-alkyl C).

Table 4

The mass losses during specific temperature ranges under no manure (Control), low manure (LM), high manure (HM) and high manure plus lime (HML) treatments. Exo_1 represents thermally labile soil organic matter (SOM); Exo_2 represents thermally stable SOM; Exo_{tot} represents total SOM. TG-T₅₀ indicates the temperature at which half of the total SOM is lost.

| Treatments | Exo ₁ (180~380°C) (%) | Exo ₂ (380~530°C) (%) | Exo _{tot} (180~530°C) (%) | Exo ₁ /Exo _{tot} | TG-T ₅₀ (°C) | |
|------------|--|--|------------------------------------|--------------------------------------|-------------------------|--|
| Control | 2.28c | 2.36b | 4.64b | 0.49b | 381a | |
| LM | 2.62b | 2.54a | 5.16a | 0.51a | 376ab | |
| HM | 2.69ab | 2.68a | 5.37a | 0.50b | 369b | |
| HML | 2.89a | 2.56a | 5.45a | 0.53a | 371b | |

Values are the means (n=3). Different lowercase letters after values in the same row indicate a significant difference among four manure treatments (P < 0.05).

Figure legends

717 Fig. 1 CPMAS-13C-NMR spectra under no manure (Control), low manure (LM), high manure (HM) and high 718 manure plus lime (HML) treatments. Alkyl (0-45 ppm); O-alkyl C (45-110 ppm); aromatic C (110-160 ppm) and 719 carbonyl C (160-220 ppm) 720 721 Fig. 2 Thermogravimetry (TG) curves (a) and corresponding derivative thermogravimetry (DTG) curves (b) of soil 722 organic matter (SOM) under no manure (Control), low manure (LM), high manure (HM) and high manure plus 723 lime (HML) treatments 724 725 Fig. 3 The relationships between SOC and variables related to the chemical protection and physical protection 726 727 Fig. 4 Relationship between soil pH and iron oxides concentrations. 728 729 Fig. 54 Biplots of the principal component analysis (PCA) between the quantity and quality of SOC and variables 730 related to chemical protection, physical protection across four manure application treatments (CK, Control; LM, 731 low manure; HM, high manure; and HML, high manure plus lime) 732 733 Fig. 65 The relationships between SOC quantity, chemical recalcitrance and thermal stability

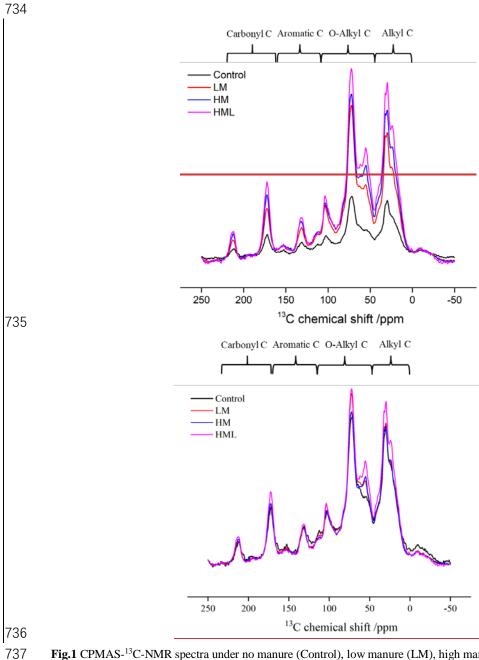


Fig.1 CPMAS-¹³C-NMR spectra under no manure (Control), low manure (LM), high manure (HM) and high manure plus lime (HML) treatments. Alkyl (0-45 ppm); O-alkyl C (45-110 ppm); aromatic C (110-160 ppm) and carbonyl C (160-220 ppm). The spectra are normalized to the total integral area of the control spectrum to highlight differences in the relative abundances of carbon functional groups.

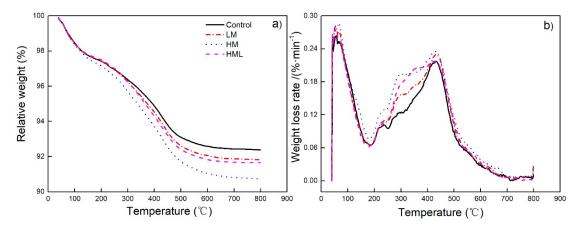


Fig. 2 Thermogravimetry (TG) curves (a) and corresponding derivative thermogravimetry (DTG) curves (b) of soil organic matter (SOM) under no manure (Control), low manure (LM), high manure (HM) and high manure plus lime (HML) treatments

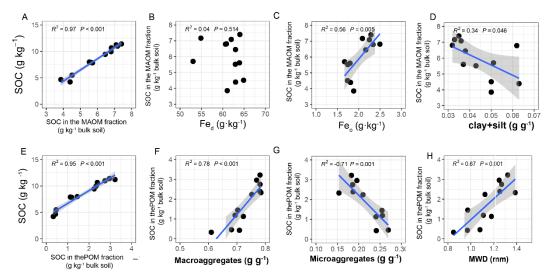


Fig. 3 The relationships between SOC and variables related to the chemical protection and physical protection. (A to D) Variables characterizing chemical protection, including the content of SOC stored in the MAOM fraction, the content of crystalline Fe oxides (Fe₀), the content of non-crystalline Fe oxides (Fe₀), and the content of clay and silt. (E to H) Variables characterizing physical protection, including the content of SOC stored in the POM fractions, the content of soil macroaggregates, the content of soil microaggregates and the mean weight diameter (MWD)

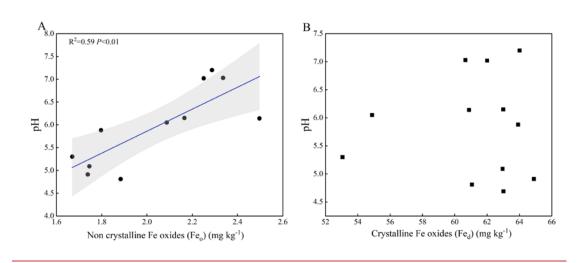


Fig. 4 Relationship between soil pH and iron oxides concentrations. (A) Significant positive correlation between pH and amorphous iron oxides (Fe₀). (B) No significant correlation between pH and crystalline iron oxides (Fe_d).

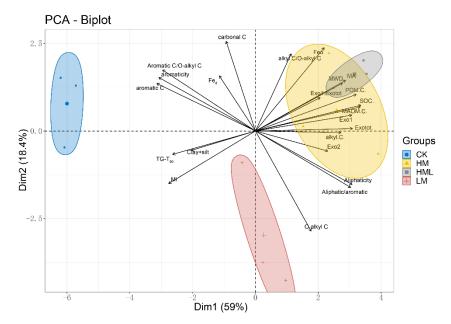


Fig. $\underline{54}$ Biplots of the principal component analysis (PCA) between the quantity and quality of SOC and variables related to chemical protection, physical protection across four manure application treatments (CK, Control; LM, low manure; HM, high manure; and HML, high manure plus lime). Fe_d, crystalline Fe oxides; Fe_o, non-crystalline Fe oxides; MA, macroaggregates (>0.25 mm); MI, microaggregates (0.05-0.25 mm); MWD, mean weight diameter; aromaticity, aromatic C / (alkyl C + O-alkyl C + aromatic C); aliphatic C, Alkyl C + O-alkyl C; aliphaticity, (Alkyl C + O-alkyl C) / (Alkyl C + O-alkyl C + Aromatic C); Exo₁, thermally labile soil organic matter (SOM); Exo₂, thermally stable SOM; Exo_{to}, total SOM; TG-T₅₀, the temperature at which half of the total SOM was lost

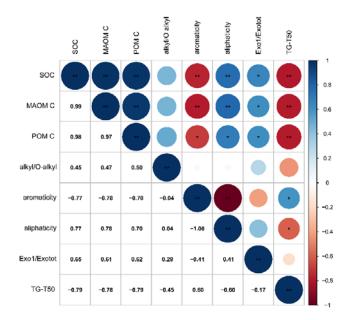


Fig. $\underline{65}$ The relationships between SOC quantity, chemical recalcitrance and thermal stability. Aromaticity, aromatic C / (alkyl C + O-alkyl C + aromatic C); aliphaticity, (alkyl C + O-alkyl C) / (alkyl C + O-alkyl C + aromatic C); Exo₁, thermally labile soil organic matter (SOM); Exo_{tot}, total SOM; TG-T₅₀, the temperature at which half of the total SOM was lost