

Dust pollution substantially weakens the impact of ammonia emission reduction on particulate nitrate formation

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18 **Abstract.** Dust emissions significantly influence air quality and contribute to nitrate aerosol pollution by altering aerosol
19 acidity. Understanding how dust interacts with ammonia emission controls is crucial for managing particulate nitrate
20 pollution, especially in urban environments. In this study, we conducted field measurements of aerosol chemical
21 components and gases across three cities in eastern China during the Spring of 2023. By combining an aerosol
22 thermodynamic model with machine learning, we assessed the relative contribution of dust to aerosol pH and its impact
23 on nitrate formation. Our results show that changes in ammonia, both in the gas and particle phases, were the main factors
24 affecting aerosol pH, with dust particles contributing to about 7% of the total pH variation. During dust events, high
25 concentrations of non-volatile ions increased aerosol pH, leading to higher nitrate levels in particle phase. Machine
26 learning analysis revealed that extreme dust storms caused a significant change in aerosol pH, enhancing nitrate
27 partitioning. Further simulations indicated that while reducing ammonia emissions is effective in lowering nitrate levels
28 under normal conditions, this effect is significantly reduced in dust-affected environments. Dust particles act as a buffer,
29 reducing the sensitivity of nitrate formation to ammonia emission reductions. These findings emphasize the need to
30 consider dust pollution when designing strategies for controlling particulate nitrate levels and highlight the complex
31 interactions between dust and anthropogenic emissions.

32

1 Introduction

Airborne dust is a major component of atmospheric aerosols, accounting for approximately 75% of the global aerosol mass load (Mahowald et al., 2006; Zhao et al., 2022; Chen et al., 2023c). Dust exerts multiple impacts on air quality (Jickells et al., 2005; Rosenfeld et al., 2001), climate (Huang et al., 2011), and human health (Zhang et al., 2023; Goudie, 2014). It can be broadly categorized into anthropogenic dust and natural dust based on sources and emission mechanisms (Chen et al., 2018; Chen et al., 2023a). Anthropogenic dust originates from human activities, such as construction, agricultural and non-exhaust vehicular emissions (Liu et al., 2021). In contrast, natural dust mainly arises from bare surfaces in arid and semi-arid regions (Shao and Dong, 2006), which cover approximately 30% of the global land area (Soussé-Villa et al., 2024; Xin et al., 2023). Beyond anthropogenic influences, over 300 countries worldwide are affected by natural dust pollution (Kurokawa and Ohara, 2020; Notaro et al., 2015). Dust storms originating in arid regions can be transported over thousands of kilometers, significantly impacting downstream air quality and atmospheric chemistry (Tan et al., 2012; Milousis et al., 2024; Sun et al., 2001).

Dust emissions contain nonvolatile cations (NVCs), such as calcium and magnesium ions, which are alkaline substances that can neutralize acidic aerosol components, such as sulfates, thereby increasing aerosol pH (Wu et al., 2013; Ding et al., 2019). Dust particles also engage in heterogeneous reactions with gaseous nitric acid, buffering acidic species and modulating pH dynamics (Zhi et al., 2025). Aerosol pH is a critical factor in atmospheric chemical processes, influencing gas-particle partitioning of inorganic aerosols (Guo et al., 2018), secondary organic aerosol (SOA) formation (Xu et al., 2015; Zhang et al., 2017; Nguyen et al., 2014), and metal-catalyzed oxidation reactions (Fang et al., 2017). Regional variations in aerosol pH alter the chemical characteristics of atmospheric pollution, affecting pollutant lifetimes and deposition rates, which in turn have profound implications for ecosystems and public health (Guo et al., 2016). Despite the incorporation of aerosol pH modules in some atmospheric chemistry models, inaccuracies in dust emission inventories can lead to biases in estimated aerosol pH, thereby introducing systematic errors in simulating associated chemical processes, such as nitrate formation.

Nitrate has emerged as a dominant component of fine particulate matter (PM_{2.5}) worldwide (e.g., China, Europe, the United States, and India), particularly as sulfate aerosol concentrations decline due to sustained SO₂ emission reductions (Weber et al., 2016; Geng et al., 2017; Zhai et al., 2021; Hauglustaine et al., 2014; Beaudor et al., 2024). The reaction between gaseous nitric acid (HNO₃) and ammonia (NH₃) represents one of the primary pathways for the formation of fine mode nitrate (Stelson and Seinfeld, 1982; Metzger et al., 2002). Nitrate formation plays a critical role in atmospheric

chemistry and the global nitrogen cycle, including reactive nitrogen deposition (Chul H. Song, 2000). The gas-particle partitioning of HNO_3 and nitrate formation is strongly influenced by aerosol pH (Guo et al., 2018; Shi et al., 2019). When total ammonia (gaseous and particulate) or NVCs are insufficient to fully neutralize aerosol sulfate, HNO_3 would not condense on aerosol due to low pH (Nenes et al., 2020; Guo et al., 2017a; Vasilakos et al., 2018; Ding et al., 2019). However, this conceptual framework may oversimplify the influence of aerosol acidity, as it fails to fully consider the substantial volatility differences between deliquescent aerosols containing sulfates or NVCs and those dominated by ammonium or nitrate, both of which are highly sensitive to aerosol pH (Nenes et al., 2020; Nenes et al., 2021). In dust-polluted environments, however, the abundance of alkaline particles, such as calcium ions, can alter nitrate formation pathways (Seinfeld et al., 1998; Hrdina et al., 2021; Li et al., 2024). Quantitative insights into how urban dust influences nitrate formation and its regulation remain nevertheless limited.

East Asia, home to some of the world's major dust source regions, significantly contributes to global atmospheric dust pollution. Under the influence of Mongolian cyclones, dust particles originating from Mongolia are transported long distances, affecting air quality and atmospheric processes across East Asia (Fu et al., 2014; Sun et al., 2001; Wang et al., 2021; Xu et al., 2020). The Yangtze River Delta (YRD) is a densely urbanized region in Eastern China, where air quality is influenced by both natural and local anthropogenic dust sources. This region provides an ideal atmospheric experiment to investigate the impact of dust pollution on urban aerosol acidity and nitrate chemistry. Under these contexts, this study examines changes in aerosol pH, and nitrate gas-particle partitioning (defined as the gas-particle partitioning of HNO_3 combined to its acid dissociation) under influence of both anthropogenic and natural dust pollution in spring 2023, focusing on three representative cities (Xuzhou, Zhenjiang, and Suzhou) in the YRD. The contributions of chemical and meteorological components to aerosol pH and the effects of dust storms on $\epsilon(\text{NO}_3^-)$ are quantified. By integrating statistical analysis approaches, we further quantify the contribution of different factors to aerosol pH and $\epsilon(\text{NO}_3^-)$. Sensitivity analyses are conducted to evaluate the effects of TNH_x ($\text{TNH}_x = \text{NH}_3 + \text{NH}_4^+$), TNO_3 ($\text{TNO}_3 = \text{HNO}_3 + \text{NO}_3^-$) and SO_4^{2-} emission controls on nitrate partitioning across varying dust pollution levels, providing a scientific basis for formulating nitrate pollution control strategies during dust events.

2.Data and Methods

2.1 Sampling site and instruments

This study selected three cities in the YRD region, China, that represent a gradient of dust transport effects: Xuzhou

(32.18°N, 119.48°E), Zhenjiang (32.16°N, 119.49°E), and Suzhou (31.29°N, 120.61°E). These cities are distributed along the north-to-south dust transport pathway, enabling a systematic investigation of the impacts of dust transport, including gradient variations in particle chemical properties, aerosol acidity (pH), and gas–particle partitioning. The sampling sites comprehensively reflect the gradient effects of dust across different regions. These sampling sites are representative of typical urban environments and reflect the general atmospheric conditions within the region.

Water-soluble inorganic ions (e.g., NH_4^+ , Na^+ , K^+ , Ca^{2+} , Mg^{2+} , SO_4^{2-} , NO_3^- , Cl^-) in $\text{PM}_{2.5}$ and gaseous components (NH_3 , HNO_3 , HCl) were continuously monitored using a Monitor for Aerosols and Gases in ambient Air (MARGA) system (Trebs et al., 2004; Rumsey et al., 2014). The system exhibited high correlation between cation and anion measurements (Fig. S1). Throughout the observation period, ambient air samples were drawn into the system, where aerosols and gaseous pollutants were separated. Water-soluble gases were removed using a wet rotating denuder, while aerosol particles with an aerodynamic diameter smaller than approximately 2.5 μm were collected using a steam jet aerosol collector (Rumsey et al., 2014; Trebs et al., 2004). Aerosol particles collected using a wet sampler were dissolved in water to form sample liquid, and then analyzed via ion chromatography. For gaseous pollutants, air samples passed through a membrane filter to remove particles before entering a scrubbing tower, where gas-phase components were dissolved in water to form sample liquid for ion chromatographic analysis (Rumsey et al., 2014). The MARGA system is equipped with automatic calibration and cleaning functions, ensuring stability and accuracy during long-term operation. The entire process is controlled by dedicated software, enabling simultaneous monitoring of multiple components and real-time data output (Schaap et al., 2004).

Meteorological parameters data (air temperature and relative humidity) were obtained from corresponding observation stations, while additional meteorological parameters were sourced from the European Centre for Medium-Range Weather Forecasts (ECMWF) ERA5 reanalysis dataset (<https://cds.climate.copernicus.eu/>, last access: November 21, 2023). Regional PM_{10} data were retrieved from the China National Environmental Monitoring Centre (<https://air.cnemc.cn:18007/>, last access: November 21, 2023).

2.2 Aerosol pH estimation

Aerosol pH is a particle property that significantly influences aerosol formation, yet it is challenging to measure directly. Traditional methods, such as ion balance and molar ratio approaches, often fail to provide accurate evaluations of aerosol pH (Guo et al., 2016; Weber et al., 2016). Currently, the most widely used approaches include the E-AIM and

ISORROPIA-II thermodynamic model (Fountoukis and Nenes, 2007), while recent studies have also begun to explore alternative methods for direct pH measurement (Li et al., 2025). In this study, we employed the ISORROPIA-II thermodynamic model to estimate aerosol pH (see Eq. 1) as well as the gas–particle partitioning of water-soluble ions, semi-volatile compounds, and water content. At low RH, aerosols are unlikely to be in a completely liquid state, and secondary organic aerosols (SOA) may affect the distribution of semi-volatile compounds due to reduced diffusion within the particles, thus influencing the predicted pH values; At high RH levels, such as $RH > 95\%$, aerosols may deliquesce, and the exponential increase in water activity (W_i) can introduce significant uncertainty into the pH values (Guo et al., 2017; Malm and Day, 2001). To improve the model's accuracy, we applied both the forward mode for metastable aerosols and excluded data with relative humidity (RH) below 35% or above 95% (Nah et al., 2018; Guo et al., 2015). The equation used to calculate aerosol pH in ISORROPIA-II is as follows (Liu et al., 2022):

$$pH = -\log_{10} \frac{1000\gamma_{H^+}C_{H^+}}{W_i} \quad (1)$$

In the Eq. (1), γ_{H^+} represents the activity coefficient of hydrogen ions, which is generally set to 1 (Liu et al., 2022). C_{H^+} denotes the hydrogen ion concentration in the aerosol aqueous phase, expressed in $\mu\text{g m}^{-3}$. W_i refers to the water content of the aerosol phase output by ISORROPIA-II (in $\mu\text{g m}^{-3}$). By incorporating these parameters, the ISORROPIA-II model provides a reliable framework for estimating aerosol pH, allowing for accurate analysis of its variation and impact under different environmental and pollution scenarios, including those influenced by dust events.

2.3 The gas–particle partitioning of nitrate

Nitrate, owing to its volatility, exists in the atmosphere in two primary forms. In the particulate phase, it predominantly appears as semi-volatile ammonium nitrate. However, where ammonia and NVCs fail to fully neutralize aerosol sulfate, the formation of semi-volatile ammonium nitrate is inhibited. Under such conditions, nitrate tends to remain in the gaseous phase as HNO_3 , which can subsequently transform into more stable coarse-mode salts, such as $\text{Ca}(\text{NO}_3)_2$, over time (Guo et al., 2017c; Vasilakos et al., 2018; Hrdina et al., 2021). Gas–particle partitioning of nitrate [$\epsilon(\text{NO}_3^-)$] defined as the ratio between particle-phase nitrate over TNO_3 serves as a key indicator of nitrate distribution between its gaseous and particulate phases. Changes in aerosol pH, influenced by varying meteorological conditions, significantly affect $\epsilon(\text{NO}_3^-)$. This study employs Eq. (2) (Guo et al., 2018; Nenes et al., 2020) to calculate theoretical values of $\epsilon(\text{NO}_3^-)$ for each observational dataset. The results enable a detailed analysis of how variations in pH across different ranges influence the gas–particle partitioning of nitrate.

$$\varepsilon(NO_3^-) = \frac{H_{HNO_3}^* W_i RT (0.987 \times 10^{-14})}{\gamma_{NO_3^-} \gamma_{H^+} 10^{-pH} + H_{HNO_3}^* W_i RT (0.987 \times 10^{-14})} \quad (2)$$

In the equation, $H_{HNO_3}^* = H_{HNO_3} K_{n1}$ ($\text{mol}^2 \text{kg}^{-2} \text{atm}^{-1}$) represent the product of the Henry's law constant and the acid dissociation constant for HNO_3 . R is the ideal gas constant ($\text{J mol}^{-1} \text{K}^{-1}$), and T is the temperature in Kelvin (K). The temperature dependence for H_{HNO_3} and K_{n1} can be found in Clegg et al. (1998). pH is calculated using Eq. (1). The factor 0.987×10^{-14} is a unit conversion factor used to convert from atm and μg to SI units. $\gamma_{NO_3^-}$ and γ_{H^+} are the activity coefficients for NO_3^- and H^+ , respectively. Activity coefficient predicted by ISORROPIA-II are $\gamma_{NO_3^-} \gamma_{H^+} = 0.28$, $\gamma_{H^+} = 1$ (Guo et al., 2018; Guo et al., 2017b; Nah et al., 2018). In the standard S-curve, pH varies within a specific range, and this relationship is influenced by the temperature dependence of the Henry's law constant and the acid dissociation constant. This model allows for a more accurate estimation of nitrate aerosol behavior under varying environmental conditions. More detailed information about inputs and outputs for the ISORROPIA-II model can be found in Tables S1 – S3.

2.4 Multi-site concentration weighted trajectory

The concentration weighted trajectory (CWT) analysis is widely used to assess the potential origins and transport pathways of air pollutants observed at receptor sites. By integrating trajectory analysis, this approach provides insights into pollutant sources and their atmospheric transport dynamics. In this study, we employed the CWT model, coupled with backward trajectories and multi-site air quality monitoring data, to investigate the potential source regions and long-range transport of the spring 2023 dust storm event observed in Xuzhou, Zhenjiang, and Suzhou. When combined with data from multiple monitoring sites, the CWT model demonstrates enhanced robustness and reliability (Boichu et al., 2019). Briefly, multi-site CWT analysis integrates pollutant concentration data from several monitoring stations with the corresponding backward trajectories to estimate the likely origins of the observed pollutants. Air pollutant concentrations are spatially allocated to grid cells traversed by air masses, weighted by the residence time within each grid cell. Compared to single-site CWT analysis, the multi-site approach offers broader spatial coverage, minimizes site-specific biases, and increases the dataset size, thereby improving the accuracy and spatial resolution of source apportionment, particularly for complex transport patterns.

In this study, 48-hour backward trajectories at 50 meters above ground level were computed using meteorological data from the Global Data Assimilation System (GDAS). The CWT analysis was conducted using the Zefir toolkit implemented in Igor Pro (Petit et al., 2017). This methodology provided a comprehensive assessment of dust transport and source attribution, facilitating a deeper understanding of dust storm dynamics in the region.

$$CWT_{ij} = \frac{\sum_{l=1}^n C_l \tau_{ij,l}}{\sum_{l=1}^n \tau_{ij,l}} \quad (3)$$

In Eq. (3), CWT_{ij} represents the weighted concentration in the grid at the i row and j column, C_l is the pollutant concentration corresponding to the l trajectory, and $\tau_{ij,l}$ is the residence time of the trajectory in the (i,j) grid. n denotes the total number of all trajectories.

2.5 Machine learning model

Aerosol pH and $\varepsilon(\text{NO}_3^-)$ exhibit nonlinear responses to multiple influencing factors. In this study, we employed a machine learning approach to investigate the effects of extreme dust storm conditions on aerosol pH and $\varepsilon(\text{NO}_3^-)$. Specifically, we used the random forest (RF) algorithm to construct regression models tailored to aerosol pH and $\varepsilon(\text{NO}_3^-)$ for each city under investigation. The dataset for the RF regression models was divided into a training set (80%) and a test set (20%). The training set was utilized to build the models, while the test set was used to validate their performance. The input predictive features for both aerosol pH and $\varepsilon(\text{NO}_3^-)$ models included the water-soluble inorganic chemical composition of aerosols (Na^+ , SO_4^{2-} , NH_4^+ , NO_3^- , Cl^- , Ca^{2+} , K^+ , Mg^{2+}), gaseous species (NH_3 and HNO_3), and meteorological parameters (T and RH). To evaluate the model performance, we applied 5-fold cross-validation for parameter tuning. Model performance was evaluated using seven statistical metrics: Mean Absolute Error (MAE), Root Mean Squared Error (RMSE), Normalized Mean Squared Error (NMSE), Mean Bias (MB), Normalized Mean Bias (NMB), Index of Agreement (IOA), and the correlation coefficient (R). Detailed definitions and calculations for these metrics are provided in Supplementary Text 1. This machine learning based approach enabled us to quantify the complex, nonlinear relationships between aerosol properties, chemical compositions, and meteorological conditions, offering deeper insights into the drivers of aerosol pH and $\varepsilon(\text{NO}_3^-)$ under varying dust pollution scenarios.

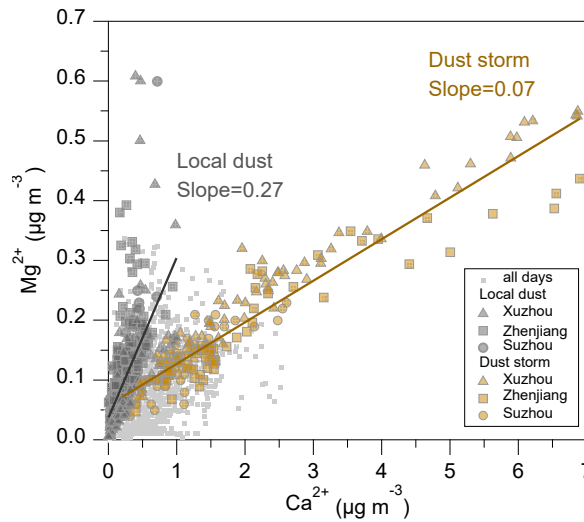
In addition, SHapley Additive exPlanations (SHAP), a method derived from the Shapley value concept in game theory, provides an interpretable framework to explain the predictions of complex machine learning models. SHAP quantifies the contribution of each input variable to individual predictions, making it a powerful tool for understanding model behavior (Duan et al., 2024; Lundberg and Lee, 2017). In this study, SHAP values were employed to assess the influence of various factors on aerosol pH and $\varepsilon(\text{NO}_3^-)$ under dust storm and local dust conditions. A positive SHAP value for a given factor indicates that it contributes positively to the prediction, whereas a negative SHAP value implies a suppressive or inhibitory impact. This analysis allowed us to disentangle the relative contributions of chemical composition, meteorological conditions, and other variables to the variations in aerosol properties under different dust

196 scenarios.

197 3. Results and Discussion

198 3.1 Observational evidence of anthropogenic and natural dust pollution

199 Dust emissions can be classified into anthropogenic and natural sources, with Ca^{2+} and Mg^{2+} commonly used as
200 tracers. Figure 1 shows the relationship between the concentrations of Ca^{2+} and Mg^{2+} during the observation period from
201 March to April 2023 across the three cities (Xuzhou, Zhenjiang, and Suzhou). It is evident that the concentrations of Ca^{2+}
202 and Mg^{2+} exhibit two distinctly different linear slopes, indicating that the different dust origins during this period were
203 influenced by both long-range transport dust storms and local dust emissions. In particular, during the period from April
204 11th to 13th, a severe dust storm originating was transmitted from northern regions, first impacting Hohhot, and then
205 southward to the southern cities of the YRD region. As shown in Fig. 2a, the PM_{10} concentrations in the cities along the
206 transport path exhibited a distinct gradient, with peak values reaching approximately $1702 \mu\text{g m}^{-3}$ in Hohhot, $1614 \mu\text{g}$
207 m^{-3} in Xuzhou, $925 \mu\text{g m}^{-3}$ in Zhenjiang, and $576 \mu\text{g m}^{-3}$ in Suzhou, respectively. In Xuzhou, the average concentration
208 of Ca^{2+} increased from $0.47 \pm 0.36 \mu\text{g m}^{-3}$ during the local dust period to $2.00 \pm 1.66 \mu\text{g m}^{-3}$ during the dust storm period,
209 marking a fourfold increase. Similarly, the average Ca^{2+} concentration rose from $0.30 \pm 0.23 \mu\text{g m}^{-3}$ to $1.69 \pm 1.41 \mu\text{g m}^{-3}$
210 in Zhenjiang, while the concentration increased from $0.35 \pm 0.26 \mu\text{g m}^{-3}$ to $0.92 \pm 0.52 \mu\text{g m}^{-3}$ in Suzhou.



211
212 **Figure 1.** Relationship between Ca^{2+} and Mg^{2+} concentrations in $\text{PM}_{2.5}$ in Xuzhou (triangle), Zhenjiang (square), and Suzhou (circle).
213 Dust types are distinguished based on the slope of the Ca^{2+} to Mg^{2+} concentration ratio, with local dust (gray) and dust storm (brown)
214 indicated. Light gray dots represent the concentrations of Ca^{2+} and Mg^{2+} observed in the three cities during March – April 2023.

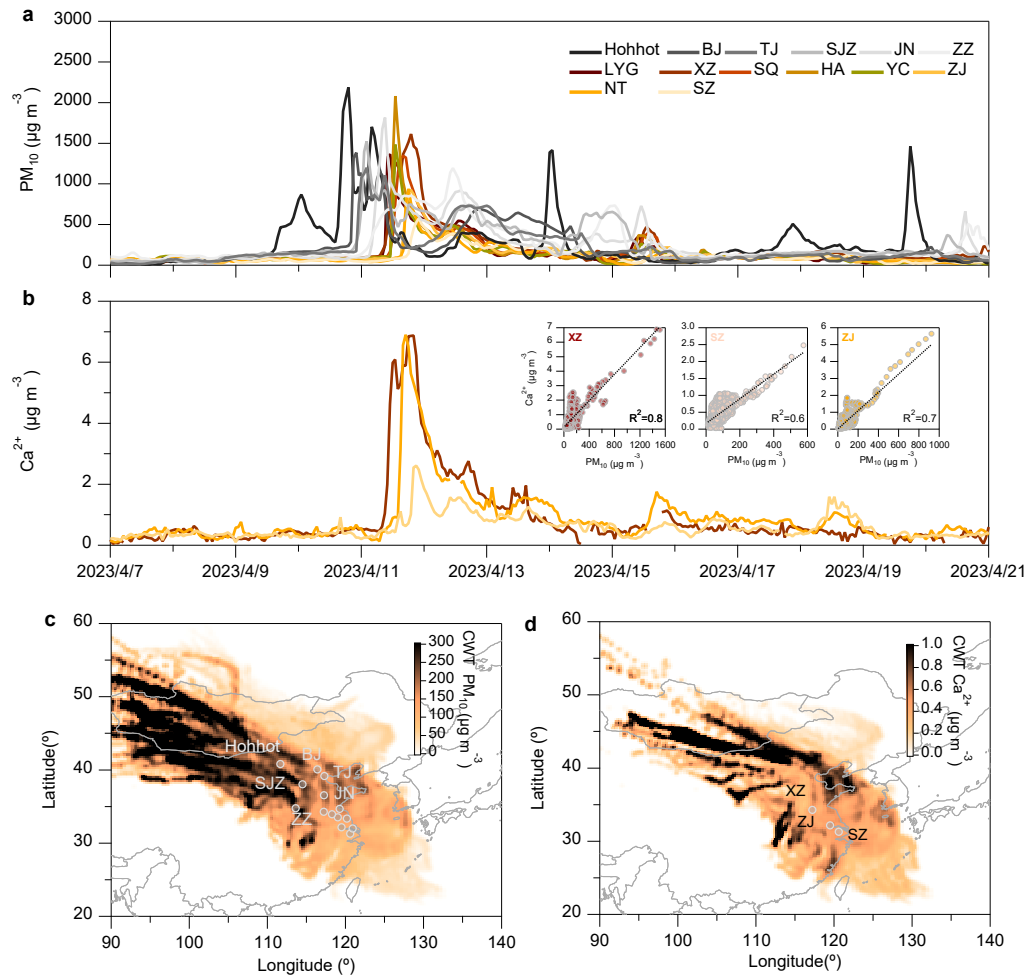


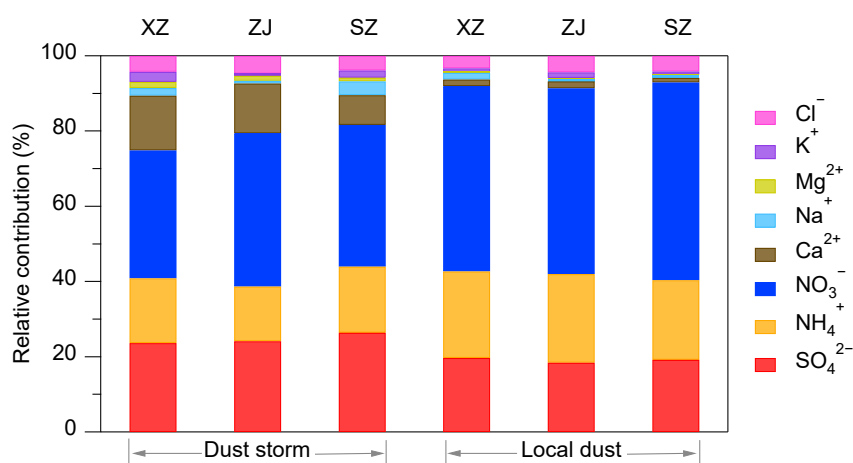
Figure 2. Time series of PM₁₀ and Ca²⁺ concentrations, and their concentration-weighted trajectories for cities along the dust transport path. **(a)** Time series of PM₁₀ in 14 cities along the BTH region, and **(b)** Time series of Ca²⁺ concentrations in Xuzhou, Zhenjiang, and Suzhou and the correlation of Ca²⁺ and PM₁₀. **(c)** 48-hour CWT-weighted spatial distribution of PM₁₀ concentrations in 14 cities from April 5 to 20, and **(d)** 48-hour CWT-weighted spatial distribution of Ca²⁺ concentrations in Xuzhou, Zhenjiang, and Suzhou (units: μg m⁻³).

Figure 2a and b illustrate the temporal evolution of PM₁₀ and Ca²⁺ concentrations during the dust storm, showing an initial spike in Hohhot, followed by a gradual increase across the Beijing-Tianjin-Hebei (BTH) region, and eventual dispersion into several cities in Jiangsu Province. This progression is consistent with the CWT-weighted trajectory patterns shown in Fig. 2c and d, which delineate the transport pathways of the dust storm. The maps highlight significant contributions from Mongolia – the dust storm's origin – to regions including Hohhot, Beijing, Tianjin, Shijiazhuang, Jinan, Zhengzhou, and Jiangsu. This finding corroborates the results of Chen et al. (2023b), who attributed the dust storm to a strong cold high-pressure system and cold front that transported substantial quantities of coarse dust aerosols southward

229 into the YRD region. Southward-moving cold fronts play a critical role in the diffusion and transport of atmospheric
230 pollutants. In arid and semi-arid regions, these storms mobilize large amounts of crustal elements, such as Ca^{2+} , with high
231 winds lifting dust from surface sources, including city streets, construction sites, and other exposed land areas (Ding et
232 al., 2019).

233 Figure 3 presents the relative contributions of water-soluble inorganic species (WSIS) in $\text{PM}_{2.5}$ during local dust and
234 dust storm periods in Xuzhou, Zhenjiang, and Suzhou. Across all three cities, the combined contribution of sulfate, nitrate,
235 and ammonium consistently accounted for over 80% of WSIS, highlighting the important role of secondary inorganic
236 aerosols in fine particulate pollution. Among these, nitrate was the most abundant species during both periods, particularly
237 during local dust events, with an average contribution ranging from 49.3% to 52.6%. However, during dust storms, its
238 relative contribution declined to 34.0 – 40.8%. In contrast, the relative contribution of sulfate increased, with increments
239 of 5.2%, 5.0%, and 6.7% observed in Xuzhou, Zhenjiang, and Suzhou, respectively. Similar trends in the relative increase
240 of sulfate and reduction of nitrate during dust events have also been widely reported in previous studies (e.g., Song et al.,
241 2023; Zhu et al., 2022). This shift in WSIS composition was likely influenced by both meteorological conditions and
242 chemical processes associated with dust. To evaluate this hypothesis, we constructed separate random forest regression
243 models for nitrate and sulfate concentrations, using meteorological parameters (Table S4) as input predictors. SHAP
244 analysis was then applied to quantify the aggregated contributions of dispersion-related variables, such as winds and
245 planetary boundary layer height. As shown in Fig. S2, nitrate exhibited a stronger response to dispersion and dilution
246 effects than sulfate, indicating its higher sensitivity to meteorological variability during the dust storm. In addition to
247 meteorological effects, heterogeneous reactions involving mineral dust also likely influence the observed variations in
248 WSIS. For instance, $\text{Ca}(\text{NO}_3)_2$ and $\text{Mg}(\text{NO}_3)_2$ coatings could tend to form preferentially on aged mineral particles rich in
249 calcite and dolomite. Previous studies have shown that the abundance of such nitrate-coated particles increases with dust
250 transport distance due to their relatively low deliquescence relative humidities ($\text{DRH} > 11\%$), which facilitate nitric acid
251 uptake under humid conditions (Li and Shao, 2009; Tobo et al., 2010; Laskin et al., 2005). Given the widespread presence
252 of calcite and dolomite in Asian dust, long-range transported particles during dust storms provide abundant alkaline
253 surfaces for heterogeneous nitrate formation. Consistent with this, our results showed an average increase of
254 approximately 10% in the relative contributions of Ca^{2+} and Mg^{2+} during the dust storm period across all three cities
255 compared to local dust events. This enhancement in alkaline mineral content suggests more effective neutralization of
256 acidic species such as HNO_3 and H_2SO_4 , thereby promoting the formation of particulate nitrates and sulfates during

257 regional dust transport. While secondary chemical formation was possible, meteorological dispersion and dilution appear
 258 to be the dominant factors leading to the observed concentration decreases.



259
 260 **Figure 3.** Relative contributions of water-soluble inorganics (SO₄²⁻, NH₄⁺, NO₃⁻, Ca²⁺, Na⁺, Mg²⁺, K⁺, and Cl⁻) within the PM_{2.5}
 261 fraction in Xuzhou, Zhenjiang, and Suzhou during dust storm and local dust pollution periods, respectively.

262

263 3.2 Driving factors of aerosol pH

264 Aerosol pH plays a crucial role in influencing aerosol formation and chemical composition. By regulating the
 265 partitioning of semi-volatile compounds between the gas and particle phases, aerosol pH directly affects the distribution
 266 of particulate matter in the atmosphere (Guo et al., 2017b). To examine the factors influencing aerosol pH, we utilized
 267 the ISORROPIA-II thermodynamic model and sensitivity analysis to quantify the relative contributions of chemical and
 268 meteorological factors, such as T and RH, in Xuzhou, Zhenjiang, and Suzhou. The correlation between simulated and
 269 observed concentrations of NH₃ and particulate NO₃⁻ is presented in Fig. 4. Across all three cities, the simulated values
 270 exhibit strong agreement with measurements ($R^2 = 0.94 - 0.99$). Additionally, Fig. S3 shows high correlations ($R^2 = 0.90$
 271 $- 0.97$) for particle-phase ammonium and chloride between ISORROPIA-II predictions and observations, confirming the
 272 robust performance of the thermodynamic model in this study.

273 To assess the impact of individual factors (TNO₃, TNH_x, Ca²⁺, SO₄²⁻, T and RH) on aerosol pH, we estimated their
 274 relative contributions using methods like those proposed by Zheng et al. (2020) and Zheng et al. (2022). First, we
 275 calculated the monthly average values for each factor in March and April, referred to as $pH_{i(3,3)}$ and $pH_{i(4,4)}$, respectively.
 276 Here, pH_i represents the influence of factor i on pH, with the numbers in parentheses indicating the respective months.
 277 For example, for the analysis of a specific factor, we used the March average value of that factor while holding the other

variables at their average levels for April. This yielded the aerosol pH value, denoted as $pH_{i(3,4)}$. Similarly, when using the April average value of the factor and maintaining the other variables at their March average levels, we recorded the resulting pH as $pH_{i(4,3)}$. The relative change in pH, denoted as $\Delta pH_{i(3)}$ and $\Delta pH_{i(4)}$ was calculated as the mean difference between $pH_{i(3,3)}$ and $pH_{i(4,3)}$, and between $pH_{i(4,4)}$ and $pH_{i(3,4)}$, respectively (see Eqs. 4 and 5). Finally, the overall impact of each factor on aerosol pH could be estimated (see Eq. 6).

$$\Delta pH_{i(3)} = pH_{i(3,3)} - pH_{i(4,3)} \quad (4)$$

$$\Delta pH_{i(4)} = pH_{i(4,4)} - pH_{i(3,4)} \quad (5)$$

$$\Delta pH_i = \frac{[\Delta pH_{i(3)}] + [\Delta pH_{i(4)}]}{2} \quad (6)$$

The impact of each factor could be positive or negative, which was detailed in Fig. S4. As shown in Fig. 5, atmospheric total ammonia emerged as the most significant driver of aerosol pH changes in all three cities, contributing 42%, 57%, and 43% of the observed pH in Xuzhou, Zhenjiang, and Suzhou, respectively. Total ammonia led to ΔpH_{TNH_x} increases of 0.6, 1.3, and 0.5 units in these cities during spring 2023. For Zhenjiang, T and Ca^{2+} were the next most influential factors, contributing 0.6 and 0.15 units to ΔpH_T and $\Delta pH_{Ca^{2+}}$, respectively. Sulfate exhibited the smallest influence on aerosol pH, where a concentration change of $0.3 \mu g m^{-3}$ corresponded to a $\Delta pH_{SO_4^{2-}}$ of approximately 0.05 units. These results align with the findings of Weber et al. (2016), which suggest that aerosol pH is less sensitive to changes in sulfate concentrations compared to ammonia levels.

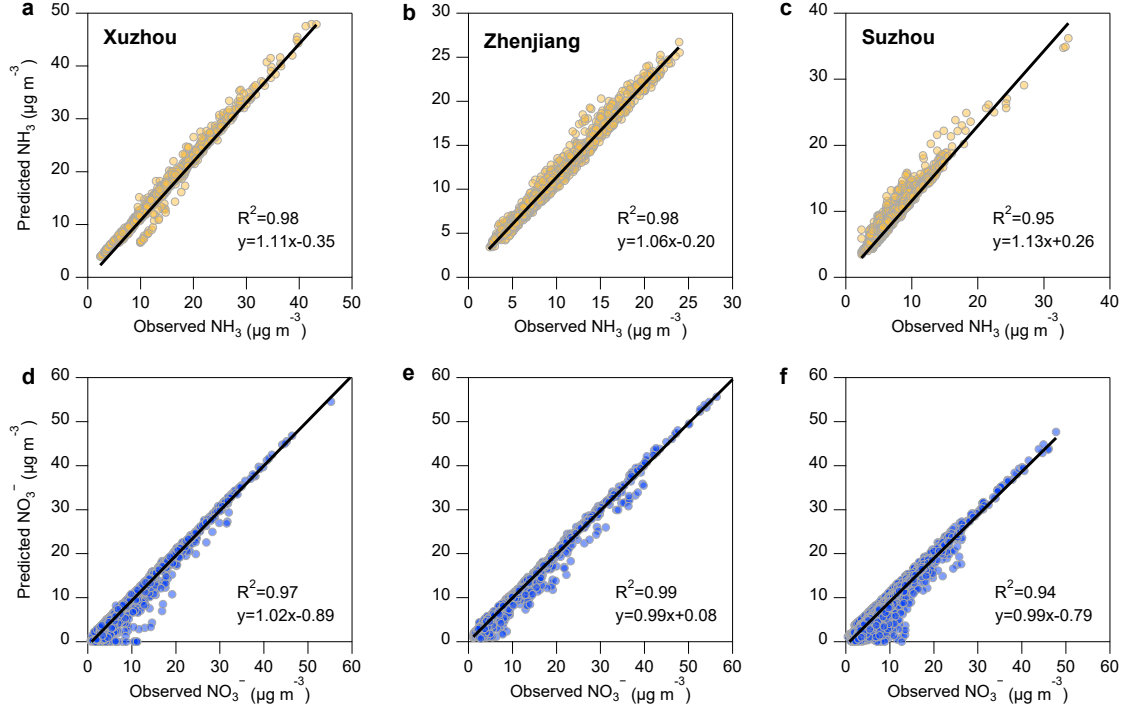


Figure 4. Correlation between ISORROPIA-II simulated and observed values of NH_3 and NO_3^- in three cities. (a) – (c) show the correlation between NH_3 predictions and observations, while (d) – (f) show the correlation between NO_3^- predictions and observations. The first column represents Xuzhou, the second column represents Zhenjiang, and the third column represents Suzhou.

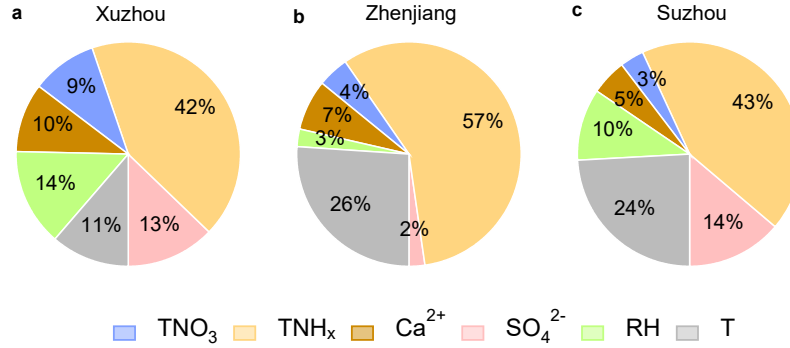
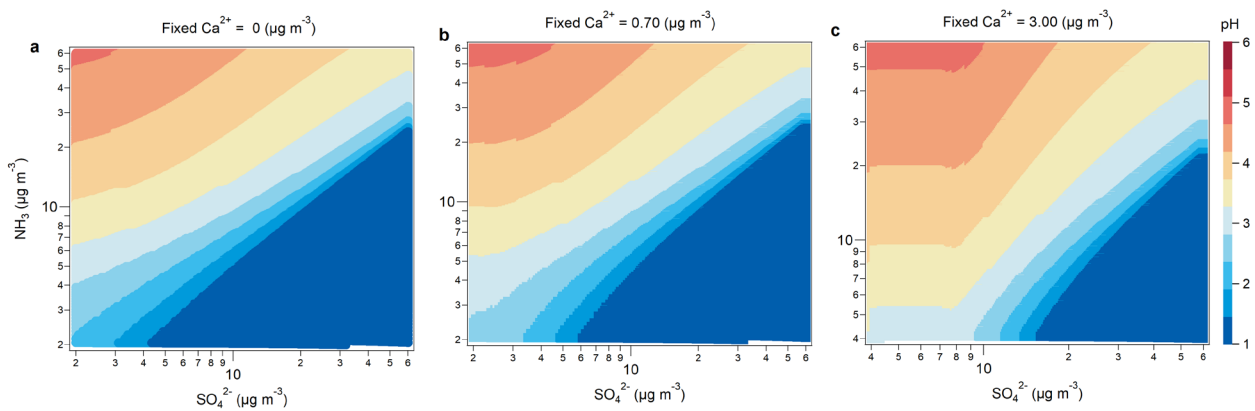


Figure 5. Relative contribution of different factors, including $\text{TNO}_3 = \text{HNO}_3 + \text{NO}_3^-$, $\text{TNH}_x = \text{NH}_3 + \text{NH}_4^+$, Ca^{2+} , SO_4^{2-} , RH, and T, to aerosol pH during the entire observation period in (a) Xuzhou, (b) Zhenjiang, and (c) Suzhou.

To further explore the response of aerosol pH to variations in SO_4^{2-} and NH_3 concentrations under different dust conditions (non-dust, local dust, and extremely dust storm), we conducted sensitivity simulations constrained by observations from Zhenjiang as a case study. As illustrated in Fig. 6a – c, we extended the NH_3 and SO_4^{2-} concentration ranges beyond their observed values to encompass potential variations across the YRD region. The input concentrations of Na^+ , SO_4^{2-} , total chloride ($\text{TCl}_x = \text{Cl}^- + \text{HCl}$), K^+ , and Mg^{2+} were fixed at the average levels observed in Zhenjiang

307 during the study period (see Table S2). Simulations were carried out under three distinct Ca^{2+} concentration scenarios: (1)
 308 non-dust ($\text{Ca}^{2+} = 0 \mu\text{g m}^{-3}$), (2) local dust ($\text{Ca}^{2+} = 0.7 \mu\text{g m}^{-3}$), and (3) extremely dust storm ($\text{Ca}^{2+} = 3.00 \mu\text{g m}^{-3}$). In these
 309 simulations, total ammonia ($\text{TNH}_x = \text{NH}_4^+ + \text{NH}_3$) and total nitrate ($\text{TNO}_3 = \text{NO}_3^- + \text{HNO}_3$) concentrations were
 310 independently changed and input into the ISORROPIA-II model. Under non-dust conditions ($\text{Ca}^{2+} = 0 \mu\text{g m}^{-3}$), the model
 311 predicted lower aerosol pH values. As shown in Fig. 6a – b, a 5 – 10-fold increase in NH_3 concentration led to a pH
 312 increase of approximately 1 unit, whereas aerosol pH demonstrated limited sensitivity to SO_4^{2-} concentration changes.
 313 This finding is consistent with previous studies (Zheng et al., 2022; Weber et al., 2016; Xie et al., 2020). However, under
 314 high Ca^{2+} concentration conditions, such as during extremely dust storm events, the influence of NH_3 on aerosol pH was
 315 notably mitigated (Fig. 6c). At relatively low SO_4^{2-} concentrations (i.e., below approximately $8 \mu\text{g m}^{-3}$, as indicated in
 316 Fig. 6c), aerosol pH exhibited diminished sensitivity to SO_4^{2-} levels, while showing greater responsiveness to variations
 317 in NH_3 . This behavior is modulated by the buffering capacity of Ca^{2+} , which preferentially reacts with SO_4^{2-} before
 318 interacting with NH_3 (Vasilakos et al., 2018), thereby limiting sulfate's ability to regulate aerosol acidity. These findings
 319 highlight that Ca^{2+} , a prominent component of mineral dust, plays a critical buffering role in mitigating the influence of
 320 NH_3 and SO_4^{2-} on aerosol acidity under dust-influenced atmospheric environments.



321 **Figure 6.** Sensitivity of the pH to ammonia (NH_3) and sulfate (SO_4^{2-}) concentrations based on ISORROPIA-II model predictions under
 322 different Ca^{2+} concentration conditions: (a) $0 \mu\text{g m}^{-3}$, (b) $0.70 \mu\text{g m}^{-3}$, and (c) $3.00 \mu\text{g m}^{-3}$.
 323

324

325 3.3 Impact of aerosol pH on the partitioning of nitric acid

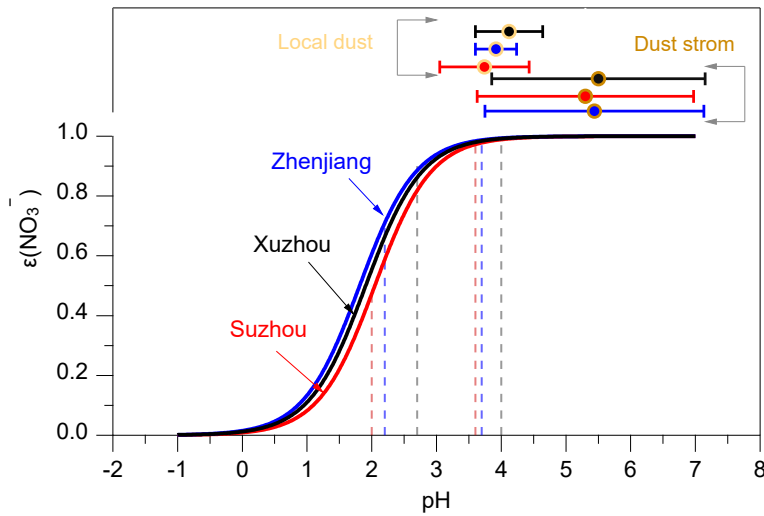
326 In eastern China, nitrate has become a key component of $\text{PM}_{2.5}$, instead of sulfate (Xu et al., 2023; Gao et al., 2023).
 327 As a semi-volatile compound, nitrate is strongly influenced by the gas-particle partitioning process in the atmosphere.
 328 Aerosol pH not only determines the stability of nitrate but also governs whether it exists in the particulate phase or

329 volatilizes as HNO_3 in the gas phase (Guo et al., 2018). At higher pH, nitrate tends to exist in the particle phase due to the
 330 oxidation of NO_x , while under lower pH conditions, nitrate is more likely to volatilize into the gas phase as HNO_3 (Nenes
 331 et al., 2020). Using Eq. (2), we analyzed the relationship between the nitrate particle-phase fraction ($\varepsilon(\text{NO}_3^-)$) and aerosol
 332 pH for three cities – Xuzhou, Zhenjiang, and Suzhou – under dust storm and local dust conditions. Fig. 7 shows the S-
 333 shaped curve representing this relationship, calculated based on the average T and aerosol W_i during dust storm and local
 334 dust conditions, assuming ideal solution behavior (activity coefficient $\gamma_{H^+} = 1$). This curve visually demonstrates the
 335 regulation of nitrate phase partitioning by aerosol pH under these conditions and provides a theoretical basis for
 336 controlling the effect of ammonia on particulate nitrate formation by adjusting aerosol pH (Guo et al., 2018).

337 As cities along the dust storm transport path, Xuzhou, Zhenjiang, and Suzhou experience varying degrees of dust
 338 influence, leading to significant differences in aerosol pH. On average, aerosol pH is elevated during dust storms
 339 compared to local dust conditions. During non-dust periods, aerosol pH values in the three cities were significantly lower
 340 than during dust events (Xuzhou: 2.7–4.0, Zhenjiang: 2.2–3.7, Suzhou: 2.0–3.6). This lower pH corresponds to a marked
 341 decrease in $\varepsilon(\text{NO}_3^-)$, indicating a shift toward gaseous HNO_3 , especially in Suzhou where $\varepsilon(\text{NO}_3^-)$ dropped to
 342 approximately 40% under the lowest pH conditions. During dust storms, the mean aerosol pH values were 5.50 ± 1.65 in
 343 Xuzhou, 5.44 ± 1.69 in Zhenjiang, and 5.30 ± 1.67 in Suzhou. Under local dust conditions, these values were lower, at
 344 4.12 ± 0.52 , 3.92 ± 0.32 , and 3.74 ± 0.69 respectively. Xuzhou, situated at the northern edge of the dust storm transport
 345 path, exhibited the highest aerosol pH during both periods, reflecting the substantial impact of transported dust pollution.
 346 The S-shaped curve in Fig. 7 demonstrates that under both dust storm and local dust conditions, the average aerosol pH
 347 aligns with nitrate particle-phase fractions exceeding 99%, indicating that nitrate predominantly resides in the particle
 348 phase. This finding highlights the promoting effect of dust pollution on the gas-to-particle transformation of nitrate.

349 When aerosol pH drops below 3, however, $\varepsilon(\text{NO}_3^-)$ decreases sharply, signifying the onset of nitrate volatilization
 350 into the gas phase. Notably, when aerosol pH lies in the range of 1 to 3, $\varepsilon(\text{NO}_3^-)$ exhibits heightened sensitivity to aerosol
 351 pH changes. This trend was consistently observed across all three cities. Reducing NH_3 concentrations is particularly
 352 effective in influencing nitrate gas-particle partitioning when aerosol pH is within this sensitive range, offering a
 353 promising strategy to mitigate regional particulate nitrate pollution. However, environments with dust pollution may
 354 disrupt this theoretical relationship. NVCs (such as Ca^{2+}) in dust can neutralize acidic aerosol components, maintaining
 355 aerosol pH at relatively high levels (e.g., $\text{pH} > \text{approximately } 3.5$) (Fig. 7). This neutralization effect limits the ability to
 356 lower particulate nitrate concentrations solely by reducing NH_3 emissions, necessitating alternative approaches to address

357 nitrate-driven air quality challenges in dust-influenced regions.



358
 359 **Figure 7.** S-curve distributions for $\epsilon(\text{NO}_3^-)$ under the conditions from different cities. Based on Eq. (2), the relationship between $\epsilon(\text{NO}_3^-)$
 360 and pH was calculated using the average T and W_i during dust storm, local dust and non-dust periods (assuming $\gamma_{\text{NO}_3^-}\gamma_{\text{H}^+}=0.28$, $\gamma_{\text{H}^+}=1$). The vertical dashed lines represent the minimum (left side) and maximum (right side) pH values under local-dust conditions
 361 calculated using ISORROPIA-II for the three cities. Error bars indicate the sample standard deviation of aerosol pH during local dust
 362 and dust storm events.
 363

364

365 To further quantify the impact of dust storms on aerosol pH and $\epsilon(\text{NO}_3^-)$, we utilized the RF model combined with
 366 SHAP values for both prediction and sensitivity analysis. The correlation between the observed and predicted results from
 367 the RF model is shown in Fig. S5. The Index of Agreement (IOA) values ranged from 0.93 to 0.97, indicating a high level
 368 of model agreement. Meanwhile, the correlation coefficients (R) varied between 0.78 and 0.90, further validating the
 369 model's predictive accuracy. For aerosol pH predictions, five evaluation metrics were used: MAE, RMSE, NMSE, MB,
 370 and NMB. The values for MAE ranged from 0.13 to 0.18, while RMSE values were between 0.26 and 0.29. For NMSE,
 371 the values ranged from 0.10 to 0.12, and the biases (MB and NMB) varied from -0.01 to -0.006 and 0.004 to 0.007,
 372 respectively. In comparison, the corresponding evaluation metrics for $\epsilon(\text{NO}_3^-)$ were as follows: MAE ranged from 0.01
 373 to 0.02, RMSE from 0.03 to 0.04, and NMSE from 0.10 to 0.21. The bias values for $\epsilon(\text{NO}_3^-)$ ranged from -0.00006 to
 374 0.004 for MB and from 0.003 to 0.007 for NMB. These statistical results demonstrate the reliability and robustness of the
 375 RF model in predicting aerosol pH and nitrate partitioning.

376 Figure 8 illustrates the impact of dust storms and local dust conditions on aerosol pH and $\epsilon(\text{NO}_3^-)$. The ΔSHAP
 377 values represent the difference between the average SHAP values of all variables during dust storm periods and the
 378 average SHAP values for all variables during the non-dust storm period. During dust storm conditions, ΔSHAP

significantly increased in Xuzhou, Zhenjiang, and Suzhou, with aerosol pH values rising by $\Delta 1.2$, $\Delta 1.5$, and $\Delta 1.5$ units, respectively (Fig. 8 a-c). This result is consistent with our previous conclusion that dust storms contribute to an increase in aerosol pH, confirming the positive impact of dust storms on the random forest model's predictions of aerosol pH. Similarly, Fig. 8 d-f shows the changes in $\varepsilon(\text{NO}_3^-)$ for the three cities under different weather conditions. It is evident that the effect of dust storms on $\varepsilon(\text{NO}_3^-)$ is 10 to 20 times greater than the impact of local non-dust storm conditions, likely due to differences in aerosol composition and enhanced alkaline inputs such as Ca^{2+} . This indicates that dust storm conditions have a significantly stronger positive contribution to the particle-phase fraction of nitrate. The presence of dust particles facilitates the conversion of nitrate to the particulate phase, highlighting the significant influence of dust storms on nitrate partitioning in the atmosphere.

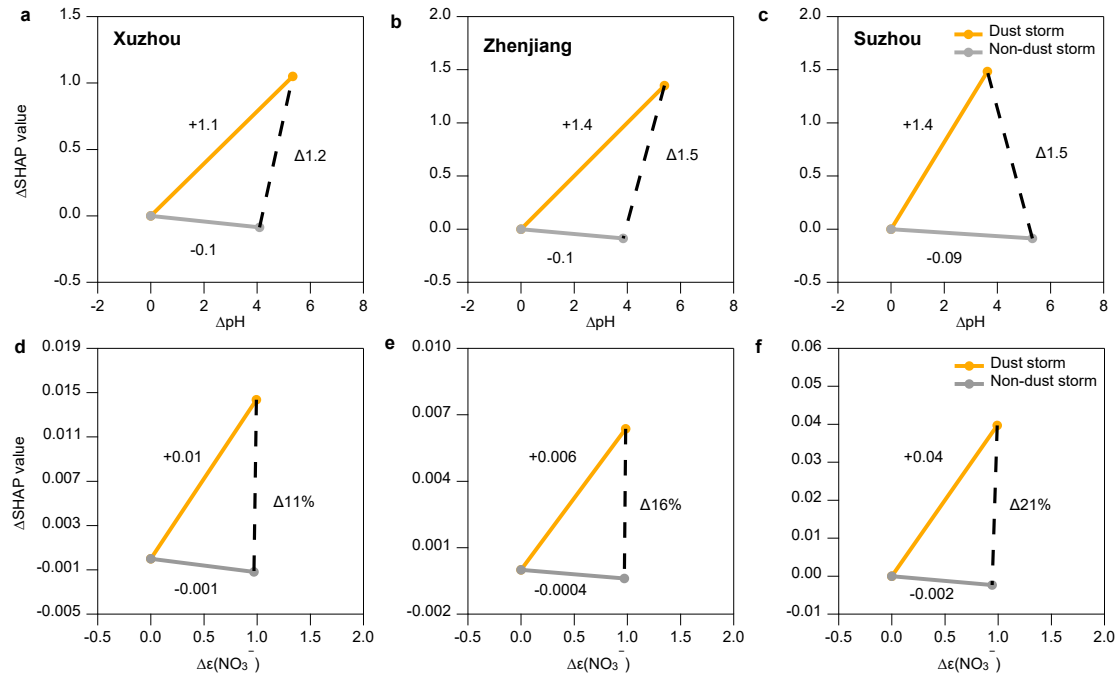


Figure 8. ΔSHAP values for (a) – (c) aerosol pH and (d) – (f) $\varepsilon(\text{NO}_3^-)$. The orange solid line represents the impact of dust storms, the gray solid line represents the non-dust scenario, and the black dashed line shows the difference between the two scenarios.

3.4 Effectiveness of emission reduction on particulate nitrate under dust pollution

To explore the impact of emission reductions of TNH_x , TNO_3 , and SO_4^{2-} on $\varepsilon(\text{NO}_3^-)$ during different dust storm conditions, we conducted a sensitivity analysis based on the thermodynamic model ISORROPIA-II, using the average pollutant concentrations observed in Zhenjiang during the spring of 2023. The results, shown in Fig. 9, demonstrate a nonlinear response of both $\varepsilon(\text{NO}_3^-)$ and the total ammonium-nitrate concentration ($\text{NH}_4^+ + \text{NO}_3^-$) to reductions in TNH_x ,

397 TNO₃, and SO₄²⁻, respectively. We simulated the effects of progressively reducing TNH_x, TNO₃, and SO₄²⁻ by 0% to 50%
398 under different Ca²⁺ concentration conditions, which include different dust pollution scenarios. For the simulation, Ca²⁺
399 concentration was set to 0.1 to 0.7 μg m⁻³ for local dust conditions and ranged from 1.0 to 3.0 μg m⁻³ for dust storm
400 conditions. When the Ca²⁺ concentration exceeded 3 μg m⁻³, further reductions in the other variables had negligible effects
401 on the output, with emission reductions having little to no impact on ε(NO₃⁻).

402 As shown in Fig. 9a, it is evident that during local dust conditions, ε(NO₃⁻) remained relatively constant until TNH_x
403 emissions were reduced by 30%. At this point, ε(NO₃⁻) rapidly dropped from 99%, signaling the onset of a significant
404 shift in the gas-particle partitioning of nitrate. When TNH_x reductions reached 50%, ε(NO₃⁻) fell sharply to approximately
405 30%, indicating that nitrate transitioned predominantly into its gas-phase form. This simulation result is consistent with
406 the sensitivity analysis of NH₃ concentrations in section 3.2, which also showed a significant response in nitrate
407 partitioning as NH₃ concentrations decreased. Thus, in the Zhenjiang region, a 30% reduction in TNH_x emissions is
408 necessary to effectively reduce the mass of (NH₄⁺ + NO₃⁻) during spring (Fig. 10d). In contrast, during dust storm
409 conditions (Fig. 9a), the reduction in TNH_x had a much more subdued effect on ε(NO₃⁻), especially at higher Ca²⁺
410 concentrations (above 2.5 μg m⁻³), where the reduction had almost no impact on ε(NO₃⁻).

411 For TNO₃ reductions, as shown in Fig. 10 b, the changes in ε(NO₃⁻) were minimal, regardless of the Ca²⁺
412 concentration. However, during local dust conditions (Fig. 9e), the reduction of TNO₃ led to a significant decrease in
413 (NH₄⁺ + NO₃⁻) concentrations, indicating that TNO₃ reduction was particularly effective under local dust conditions.
414 Lastly, reductions in SO₄²⁻ emissions (Fig. 9c and f) had a smaller impact on both ε(NO₃⁻) and (NH₄⁺ + NO₃⁻)
415 concentrations. Interestingly, at very low dust concentrations, SO₄²⁻ reductions could even lead to a slight increase (by up
416 to 0.5%) in ε(NO₃⁻), indicating that sulfate reduction alone is not an effective strategy for controlling nitrate partitioning.

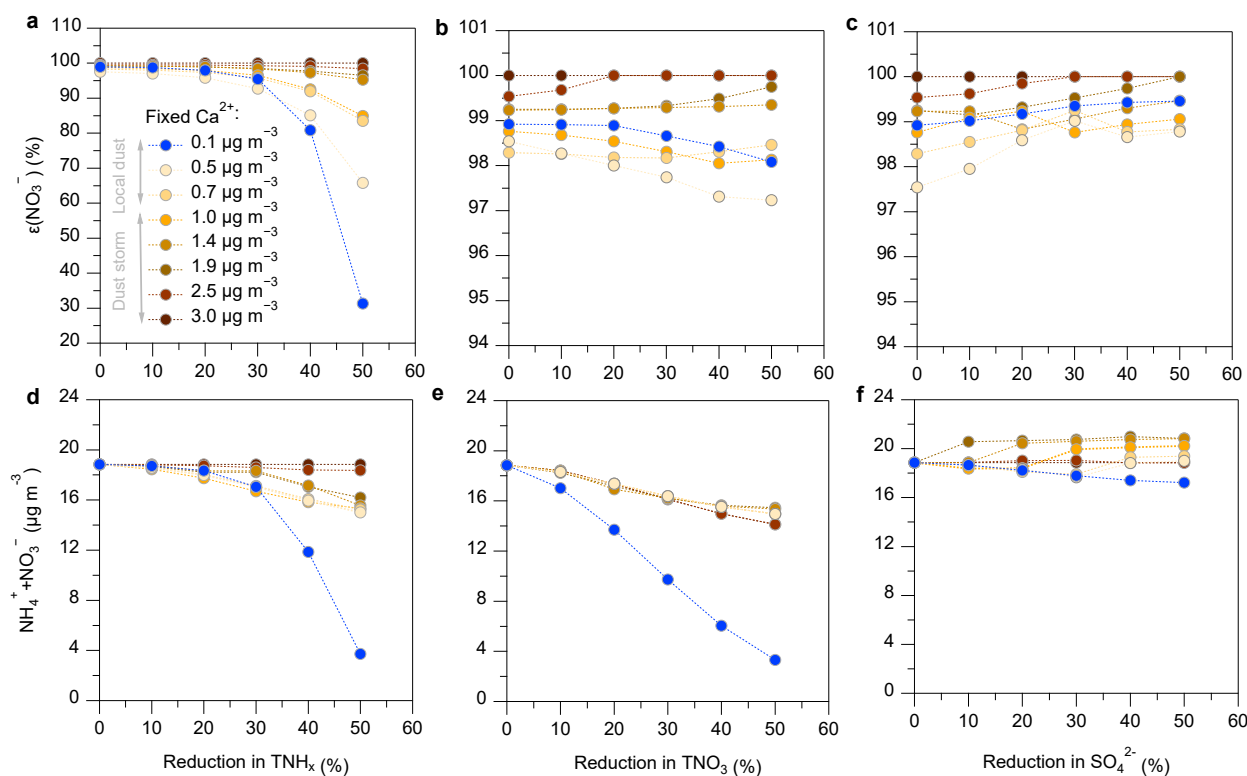


Figure 9. Sensitivity analysis based on the thermodynamic model ISORROPIA-II simulated the impact of reducing TNH_x ($\text{TNH}_x = \text{NH}_3 + \text{NH}_4^+$), TNO_3 ($\text{TNO}_3 = \text{HNO}_3 + \text{NO}_3^-$), and SO_4^{2-} by 0-50% during dust events of varying intensities on $\epsilon(\text{NO}_3^-)$ and $\text{NH}_4^+ + \text{NO}_3^-$.

4. Conclusions and Impactions

This study explores the impact of dust pollution on aerosol pH and nitrate gas-particle partitioning in three cities across the YRD region of Eastern China. By combining field observations, thermodynamic modeling, and machine learning techniques, we provide a comprehensive analysis of how different dust scenarios affect urban aerosol pH and gas-particle partitioning chemistry of nitrate. Our analysis of a dust storm event that originated in Mongolia and was transported over long distances to the YRD region in the spring of 2023 revealed a significant increase in PM_{10} concentrations, the average PM_{10} concentration in three cities along the route exceeds $400 \mu\text{g m}^{-3}$, approximately four times higher than during local dust events. Thermodynamic simulations using the ISORROPIA model showed that both ammonia and calcium ion concentrations strongly influenced aerosol pH, with average contributions of 47% and 7% respectively. Random forest model simulations further indicated that the presence of high NVCs during dust storms significantly contributed to changes in aerosol pH (1.2 – 1.5 units). Sensitivity analysis of pH responses to sulfate and NH_3 concentrations under different dust conditions (non-dust, local dust, and extremely dust storm) revealed that a 5 to 10 fold increase in NH_3 led to a 1-unit change in aerosol pH. Machine learning analysis showed that extreme dust storm

434 events contributed approximately 1.4 units to the increase in aerosol pH, with a corresponding increase in nitrate
435 partitioning (16%). This suggests that under high aerosol pH conditions during dust pollution periods, nitrate is
436 predominantly in the particulate phase, indicating that dust significantly inhibits the partitioning of nitrate into the gaseous
437 phase. In addition, our sensitivity analyses also showed that ammonia reduction had the most significant effect on reducing
438 nitrate aerosols under dust-free conditions. However, the effectiveness of ammonia reductions in lowering nitrate aerosol
439 concentrations was significantly reduced due to the influence of NVCs on nitrate partitioning under dust pollution
440 scenarios. These findings suggest that dust pollution can substantially weaken the impact of ammonia reductions on nitrate
441 aerosol formation, highlighting the need for targeted control strategies during dust storm events. Dust emission remains
442 a significant air pollution concern worldwide, while urban nitrate aerosol pollution is a pressing issue in many cities,
443 particularly in East Asia, where the frequency of natural dust events has increased in recent years. These dust storms,
444 along with anthropogenic dust, can substantially alter aerosol chemistry by modifying aerosol pH and nitrate partitioning.
445 Therefore, effective dust control strategies are critical for mitigating the adverse effects of aerosol acidity on nitrate
446 aerosol formation and improving air quality in dust-prone regions.

447

448 *Data availability.* Additional meteorological parameters can be accessed at the European Centre for Medium-Range
449 Weather Forecasts (ECMWF) ERA5 reanalysis dataset (<https://cds.climate.copernicus.eu/>; last access: 21 November
450 2023). Regional PM₁₀ data can be accessed at the China National Environmental Monitoring Centre
451 (<https://air.cnemc.cn:18007/>; 21 last access: November, 2023). The additional data will be made available upon request
452 (yjzhang@nuist.edu.cn).

453
454 *Author contributions.*
455 YZ conceived and designed the study. HL and YZ conducted the simulations and data analysis. HL, YZ, SZ, YR, JQ, and
456 MZ carried out field measurements and validated the data. HL and YZ wrote the original manuscript, while DL, FC, OF,
457 HD, and XG provided critical feedback and contributed to the manuscript revisions.

458
459 *Competing interests.* The authors declare that they have no conflict of interest.

460
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