



## 1        **Direct observation of core-shell structure and water** 2        **uptake of individual submicron urban aerosol particles**

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23       **Abstract.** Determining the particle chemical morphology is crucial for unraveling reactive uptake in  
24       atmospheric multiphase and heterogeneous chemistry. However, it remains challenging due to the  
25       complexity and inhomogeneity of aerosols particles. Using a scanning transmission X-ray microscopy  
26       (STXM) coupled with near-edge X-ray absorption fine structure (NEXAFS) spectroscopy and an  
27       environmental cell, we imaged and quantified the chemical morphology and hygroscopic behavior of  
28       individual submicron urban aerosol particles. Results show that internally mixed particles composed of  
29       organic carbon and inorganic matter (OCIn) dominated the particle population ( $73.1 \pm 7.4\%$ ). At 86%  
30       relative humidity, 41.6% of the particles took up water, with OCIn particles constituting 76.8% of these  
31       hygroscopic particles. Most particles exhibited a core-shell structure under both dry and humid  
32       conditions, with an inorganic core and an organic shell. Our findings provide direct observational

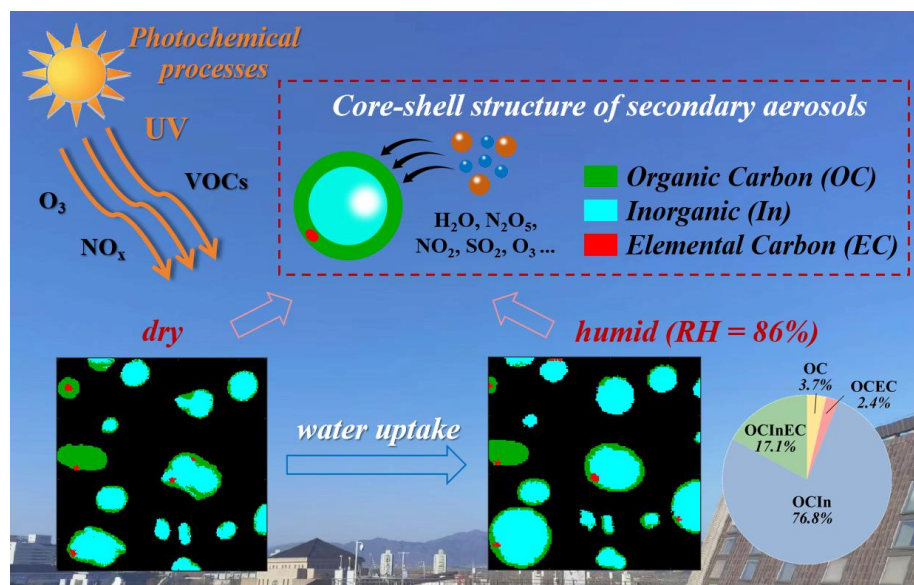


33 evidence of the core-shell structure and water uptake behavior of typical urban aerosols, which  
34 underscore the importance of incorporating the core-shell structure into models for predicting the  
35 reactive uptake coefficient of heterogeneous reactions.

36 **Short summary:** The particle chemical morphology is important to atmospheric multiphase and  
37 heterogeneous chemistry. This work directly observed the core-shell structure and water uptake  
38 behavior of individual submicron aerosol particles at an urban site and elucidated the potential impact  
39 on particle reactive uptake and heterogeneous reactions.

40 **Keywords:** urban air pollution; individual particles; chemical morphology; core-shell structure; water  
41 uptake

42 **Table of Contents Graphic:**



43



## 44 1 Introduction

45 Aerosols have significant impacts on visibility, climate, and human health (McCormick and  
46 Ludwig, 1967; Noll et al., 1968; Chow et al., 2006; Rasool and Schneider, 1971). As particles in the  
47 atmosphere usually act as reaction vessels for various reactions, the physicochemical properties of  
48 aerosol particles play an important role in reactive uptake of gaseous molecules onto particles, mass  
49 transfer and gas-particle partitioning equilibrium, and transformation mechanisms of pollutants (Abbatt  
50 et al., 2012; Davidovits et al., 2011; George et al., 2015; George and Abbatt, 2010; Su et al., 2020;  
51 Ziemann and Atkinson, 2012). Therefore, quantitatively characterizing the aerosol physicochemical  
52 properties is vital to atmospheric multiphase and heterogeneous chemistry (Freedman, 2017; Li et al.,  
53 2016; Riemer et al., 2019; Tang et al., 2016).

54 The aerosol physicochemical properties, such as the particle size, chemical morphology (defined  
55 as the spatial distribution of various chemical components within a particle herein), mixing state, and  
56 hygroscopicity vary under different ambient conditions. These properties and their variations have a  
57 critical influence on reactive uptake, a key process in multiphase chemistry which is initiated by the  
58 collision of a gas-phase reactant with a condensed-phase surface (Davies and Wilson, 2018; Reynolds  
59 and Wilson, 2025). Organic coatings of particles with core-shell structures inhibit reactive uptake of  
60 dinitrogen pentoxide ( $\text{N}_2\text{O}_5$ ) by particulate matter by means of affecting mass accommodation, the  
61 availability of water for hydrolysis, and mass transport (Wagner et al., 2013; Jahl et al., 2021; Jahn et  
62 al., 2021). Without considering the core-shell structure, the reactive uptake coefficient of  $\text{N}_2\text{O}_5$  tends to  
63 be overestimated from several times to tens of times (Wagner et al., 2013). Therefore, investigating the  
64 particle chemical morphology is necessary for accurately quantifying the uptake coefficient and  
65 reducing uncertainty in heterogeneous reactions.

66 So far, extensive research has been conducted on the physicochemical properties of bulk aerosols  
67 by various techniques, such as the Humidified Tandem Differential Mobility Analyzer (H-TDMA),  
68 Aerosol Mass Spectrometer (AMS), and Soot Particle Aerosol Mass Spectrometer (SP-AMS) (Li et al.,  
69 2016; Tang et al., 2019; Riemer et al., 2019). However, the bulk analysis mostly obtains indirect  
70 information about the physicochemical properties of particle populations based on assumptions and  
71 estimations, which is difficult to directly observe the chemical morphology and mixing state of aerosol  
72 particles (Li et al., 2016). This knowledge gap hinders our understanding of the role of particle aerosols



73 in reactive uptake and heterogeneous processes. As a comparison, individual particle analysis can  
74 provide direct observational evidence about the chemical morphology and mixing state at the  
75 microscopic scale, which is essential for exploring particle hygroscopic and optical properties (Krieger  
76 et al., 2012; Li et al., 2016; Posfar et al., 2010; Wu and Ro, 2020).

77 Scanning transmission X-ray microscopy combined with near-edge X-ray absorption fine  
78 structure (STXM/NEXAFS) spectroscopy bases on synchrotron radiation technology. It is a robust  
79 technique for obtaining chemical morphology information of numerous individual particles with high  
80 spectral energy resolution, as it can resolve compositional contrast at the single particle level within a  
81 particle population. Compared with electron microscopy which is not well-suited for analyzing organic  
82 species, the soft X-ray energy range of STXM makes it possible to quantify light elements (such as C,  
83 N, O) with high chemical specificity, high spatial resolution, and little beam damage (Moffet et al.,  
84 2011; Shao et al., 2022). In addition, STXM doesn't require ultrahigh vacuum conditions. Therefore,  
85 STXM/NEXAFS spectroscopy has enormous potential in exploring ambient samples under  
86 atmospheric relevant conditions, especially submicron-sized particles.

87 Compared with other STXM which generally analyze samples under vacuum conditions (Alpert et  
88 al., 2022; Bondy et al., 2018; Fraund et al., 2020; Knopf et al., 2023; Lata et al., 2021; Moffet et al.,  
89 2010a; Moffet et al., 2010b; Moffet et al., 2013; Moffet et al., 2016; Tomlin et al., 2022), several  
90 STXM instruments are equipped with an in-situ temperature and relative humidity (RH) control  
91 environmental cell, allowing for investigating hygroscopicity and water uptake behavior of  
92 laboratory-generated particles (Ghorai and Tivanski, 2010; O'brien et al., 2015; Piens et al., 2016;  
93 Zelenay et al., 2011a; Zelenay et al., 2011b). However, only a few researches focusing on hygroscopic  
94 behavior of ambient particles has been reported so far, and these particles were collected in rural  
95 environment (Piens et al., 2016) or forest (Mikhailov et al., 2015; Pöhlker et al., 2014). Studies on  
96 water uptake of urban aerosol particles using STXM and corresponding knowledge for their chemical  
97 morphology under humid conditions is currently lacking.

98 In this study, we investigated the chemical morphology of ambient individual submicron aerosol  
99 particles using STXM/NEXAFS spectroscopy. The ambient samples were collected at an urban site in  
100 North China Plain during a pollution episode. We also explored the chemical morphology and water  
101 uptake behavior of individual particles at high humidity (RH = 86%) using an environmental cell. This  
102 work aims to improve our comprehension of the physicochemical properties of particles in typical



103 urban pollution atmospheres, aiding in clarifying their atmospheric heterogeneous processes and  
104 multiphase chemistry.

## 105 **2 Materials and methods**

### 106 **2.1 Sampling and instruments**

107 To study the physicochemical properties of ambient particles, samples were collected during a  
108 pollution episode at the Peking University Urban Atmosphere Environment Monitoring Station  
109 (PKUERS, 39°59'21"N, 116°18'25"E) in Beijing, China. More details about the measurement site can  
110 be found in our previous studies (Tang et al., 2021; Wu et al., 2007).

111 The individual particle sample was collected using a four-stage cascade impactor with a Leland  
112 Legacy personal sample pump (Sioutas, SKC, Inc., the US) at a flow rate of 9 L min<sup>-1</sup>. The sampling  
113 started at 5:04 P.M. on October 1<sup>st</sup>, 2019 and lasted for 5 minutes. The sampling substrate was a copper  
114 grid (Lacey Carbon 200 mesh, Ted Pella, Inc., the US) suitable for its X-ray transparency. Particles  
115 collected onto the last stage with the 50% cut-point aerodynamic diameter of 250-nm were used for  
116 STXM analysis. The sample was placed into a sample box sealed with a bag filled with nitrogen, and it  
117 was stored in a freezer at a temperature of -18°C until analysis.

118 Other parameters were measured from September 28<sup>th</sup> to October 7<sup>th</sup>, 2019. The non-refractory  
119 chemical composition of submicron particles (NR-PM<sub>1</sub>) was obtained by a Long Time-of-Flight  
120 Aerosol Mass Spectrometer (LTOF-AMS, Aerodyne Research Inc., the US) (Zheng et al., 2020; Zheng  
121 et al., 2023). Calibrations of ionization efficiency (IE) and relative IE followed the standard procedures  
122 described in previous studies (Canagaratna et al., 2007; Fröhlich et al., 2013). The reference  
123 temperature and pressure conditions of mass concentrations reported herein were 293.7 K and 101.82  
124 KPa. We applied composition-dependent collection efficiency (CDCE) values ( $0.50 \pm 0.01$ , mean  $\pm$   
125 standard deviation) that were calculated by the methods introduced by Middlebrook et al. (2012) to the  
126 AMS data. The mass concentration of fine particles (PM<sub>2.5</sub>) was measured by a TEOM analyzer  
127 (TH-2000Z1, Wuhan Tianhong Environmental Protection Industry Co., Ltd., China). Meteorological  
128 parameters including temperature (*T*), RH, wind speed, and wind direction were monitored by an  
129 integrated 5-parameter Weather Station (MSO, Met One Instruments, Inc., the US).



## 130 2.2 STXM/NEXAFS analysis

131 In order to gain the chemical morphology, mixing state, and component information of individual  
132 particles, STXM/NEXAFS spectroscopy measurements were carried out at the PolLux beamline  
133 (X07DA) of the Swiss Light Source (SLS) at Paul Scherrer Institute (PSI) (Raabe et al., 2008). In brief,  
134 X-rays illuminated a Fresnel zone plate focusing the beam to a pixel of  $35 \times 35 \text{ nm}^2$ . The zone plate has  
135 a central stop that acts together with another optic known as an order sorting aperture to eliminate  
136 unfocused and higher-order light, ensuring only first-order focused light is transmitted to the sample.  
137 Then, X-rays transmitted through the sample are detected. The absorbance of each pixel is  
138 characterized by optical density (OD) based on the Beer-Lambert's law as follows,

$$\text{OD} = -\ln(I/I_0) \quad (1)$$

139 where  $I$  and  $I_0$  are the intensity of photons transmitted through a sample region and a sample-free  
140 region, respectively. Further details including the uncertainty estimation of OD are described in the  
141 Supplementary Information (SI).

142 STXM/NEXAFS spectroscopy scans X-ray energies over particles with high spectral energy  
143 resolution. When inner shell electrons of atoms absorb X-ray photons, they can transition into  
144 unoccupied valence orbitals, resulting in an absorption peak that is used to identify specific bonding  
145 characteristics. The amount of absorption depends on the photon energy ( $E$ ), elemental composition, as  
146 well as sample thickness and density (Moffet et al., 2011). We employed two measurement strategies  
147 to optimize photon flux to the particles, achieving the best signal-to-noise ratio while minimizing the  
148 scan time. The first strategy was a high energy-resolution mode with an X-ray energy resolution  $\Delta E =$   
149  $0.2 \text{ eV}$  and a coarse pixel size of around  $100 \times 100 \text{ nm}^2$  to measure absorption at small energy steps.  
150 The energy resolution is defined as being able to distinguish between two absorption peaks separated  
151 by  $\Delta E$  at the full width at half maximum OD. In this mode, carbon (C), nitrogen (N), and oxygen (O)  
152 K-edge spectra of individual particles were measured. The energy offset of C and O spectra were  $+0.4$   
153  $\text{eV}$  and  $+1.2 \text{ eV}$  respectively, according to the energy calibration procedures using polystyrene spheres  
154 and gas-phase carbon dioxide ( $\text{CO}_2$ ). The energy offset of N at the K-edge was not calibrated, however,  
155 the obtained spectra of ambient particles appeared identical to ammonium salts in literatures (Ekimova  
156 et al., 2017; Latham et al., 2017). Due to the presence of ammonium, which was confirmed in particles



157 using AMS, we applied a calibration factor of +0.1 eV for the N K-edge to match our observed main  
158 peak to that of ammonium at 405.7 eV.

159 The second strategy is a high spatial-resolution mode with a pixel size of  $35 \times 35 \text{ nm}^2$  and  $\Delta E =$   
160 0.6 eV, where imaged at four specific energies for the C K-edge, namely, 278.0 eV, 285.4 eV, 288.6  
161 eV, and 320.0 eV. Automated analysis followed the methodology of Moffet et al. (2010a). In brief,  
162 absorption at 278.0 eV ( $\text{OD}_{278.0\text{eV}}$ ) is regarded as the pre-edge of carbon, which is mainly due to  
163 off-resonance absorption by inorganic elements other than carbon. Absorption at 285.4 eV ( $\text{OD}_{285.4\text{eV}}$ )  
164 is due to the characteristic transition of  $\text{sp}^2$  hybridized carbon (i.e., doubly bonded carbon). Since this  
165 peak is abundant for elemental carbon (EC), it can be used to discern soot, because EC is a type of  
166 components of soot (Penner and Novakov, 1996). Absorption at 288.5 eV ( $\text{OD}_{288.5\text{eV}}$ ) comes from  
167 carboxylic carbonyl groups, which are common in organic aerosols in atmospheres. Therefore, organic  
168 carbon (OC) is identified by this energy. Absorption of the post-edge at 320.0 eV ( $\text{OD}_{320.0\text{eV}}$ ) is  
169 contributed by carbonaceous and non-carbonaceous atoms (Moffet et al., 2010a).

170 Based on absorption at these four typical energies, we obtain three images by further processing.  
171 The difference between OD at the post-edge and OD at the pre-edge ( $\text{OD}_{320.0\text{eV}} - \text{OD}_{278.0\text{eV}}$ ) indicates  
172 total carbon. The ratio of OD at the pre-edge to OD at the post-edge ( $\text{OD}_{278.0\text{eV}} / \text{OD}_{320.0\text{eV}}$ ) indicates the  
173 relative absorption contribution of inorganic matter (In). Compared with the absorbance contribution of  
174 doubly bonded carbon to total carbon ( $\%\text{sp}^2$ ) in the highly oriented polycrystalline graphite (HOPG,  
175 assuming that  $\%\text{sp}^2 = 100\%$ ) at 285.4 eV, the spatial distribution of EC/soot in samples can be  
176 identified by the procedure of Hopkins et al. (2007). It is assumed that total carbon consists of OC and  
177 EC. These three images described above were then overlaid to create a chemical map of individual  
178 particles.

### 179 2.3 Criterion of particle water uptake based on the total oxygen absorbance

180 To determine whether particles took up water, a criterion was established on the basis of the total  
181 oxygen absorbance determined at the energy of 525.0 eV (the pre-edge of oxygen) and 550.0 eV (the  
182 post-edge of oxygen). Based on the same principle as the total carbon calculation, the difference  
183 between OD at the post-edge and pre-edge of oxygen represents the total oxygen absorbance. Due to  
184 the fact that each particle is composed of some pixels, the total oxygen absorbance ( $\Delta\text{OD}$ ) of an  
185 individual particle under dry and humid conditions is calculated as follows,



$$\Delta OD_{dry} = \sum_{i=1}^m \Delta OD_i = \sum_{i=1}^m (OD_{post,i} - OD_{pre,i}) = \sum_{i=1}^m OD_{post,i} - \sum_{i=1}^m OD_{pre,i} \quad (2)$$

$$\Delta OD_{humid} = \sum_{j=1}^n \Delta OD_j = \sum_{j=1}^n (OD_{post,j} - OD_{pre,j}) = \sum_{j=1}^n OD_{post,j} - \sum_{j=1}^n OD_{pre,j} \quad (3)$$

where  $m$  and  $n$  are numbers of pixels that make up an individual particle under dry and humid conditions,  $i$  and  $j$  are a certain pixel within an individual particle under dry and humid conditions, *post* and *pre* respectively represent the energy at the post-edge (550.0 eV) and that at the pre-edge (525.0 eV) of oxygen. If a particle takes up water, the amount of oxygen atoms within this particle will increase, leading to an amplification in  $\Delta OD$ . Water uptake may increase particle height and absorption. On the other hand, it possibly causes a particle to spread out, which may reduce particle height and thus absorption. Although a thinner particle that contains more water may result in less absorption at some specific pixels,  $\Delta OD_{humid}$  will be larger than  $\Delta OD_{dry}$  due to the fact that more pixels are summed, i.e.,  $n > m$ . Therefore, comparing the results of Eq. 2 and Eq. 3 will quantify the total oxygen absorbance of a particle under dry and humid conditions, and determine particle water uptake. Specifically, if  $\Delta OD_{humid} > \Delta OD_{dry}$ , then we assume that the particle has taken up water.

## 2.4 A novel in-situ environmental cell

To explore the chemical morphology and hygroscopicity of the particles under humid conditions, we adjusted the RH of an in-situ environmental cell with sample placed in it. The environmental cell can also be used for trace gas reactive uptake and photochemical reactions with laboratory-generated particles (Alpert et al., 2019; Alpert et al., 2021). The environmental cell used in this study consists of a removable sample clip that hosts a sealed silicon nitride (SiNit) window and a main body that contains gas supply lines and temperature control. Together, they are mounted in the STXM vacuum chamber. A SiNit window at the back side of the main body is also sealed and ensures X-ray transparency passing through the whole environmental cell assembly. Descriptions of the connections for the gas supply, heating and cooling devices, and temperature measurement can be found in previous studies (Huthwelker et al., 2010; Zelenay et al., 2011a). The detailed methods of collecting the ambient particles by the impactor and measuring them in the environmental cell were shown in the SI.

We performed humidity calibration experiments to make sure sufficient heat transfer and a homogeneous water vapor field across the samples. It is important due to the fact that the only way for





211 samples to gain or lose heat and water was through air contact. To study the accuracy of RH in the  
212 environmental cell, water uptake and deliquescence of a sodium chloride (NaCl) standard sample was  
213 observed. The deliquescence relative humidity (DRH) of pure NaCl crystals obtained from literatures  
214 and thermodynamic models is around 75 – 76% at room temperature (Eom et al., 2014; Martin, 2000;  
215 Peng et al., 2022). The images of the NaCl sample displayed in Fig. S1 illustrate the morphological  
216 changes as RH increased. As shown in Fig. S1, particle morphologies in panels (A) – (C) remained  
217 essentially identical before RH reached the DRH of NaCl, although the focus position slightly varied in  
218 different panels. When RH was 75.6% (Fig. S1D), particles completely deliquesced and some  
219 coalesced. The uncertainty of RH in the environmental cell in this study was determined conservatively  
220 to be  $\pm 2\%$ , in agreement with previous results (Huthwelker et al., 2010). Information about the oxygen  
221 K-edge spectra of the NaCl sample at high RH can be found in Fig. S2.

## 222 **3 Results and discussion**

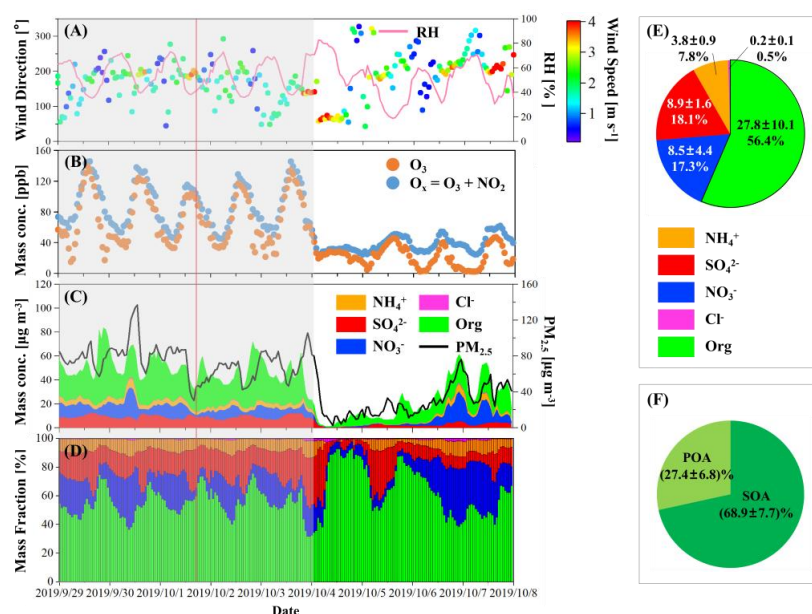
### 223 **3.1 Pollution characteristics during the sampling period**

224 Time series of meteorological parameters, mass concentrations of gaseous pollutants,  $\text{PM}_{2.5}$ , and  
225  $\text{NR-PM}_1$  are shown in Fig. 1. During the pollution episode from September 29<sup>th</sup> to October 3<sup>rd</sup>, 2019,  
226 the stagnant weather condition with low wind speed led to pollution accumulation. The air became  
227 clean due to the appearance of a strong north wind on October 4<sup>th</sup> (Fig. 1A). The sampling time of the  
228 individual particle sample was 5:04 P.M. on October 1<sup>st</sup> (see the red line in Fig. 1) with an ozone ( $\text{O}_3$ )  
229 concentration of 97.1 ppb. At that time, the low mass fraction of volatile inorganic species such as  
230 nitrate made it suitable for measurements using offline techniques, such as STXM, because the loss of  
231 volatile species during storage and measurement processes was minimal.

232 During the pollution episode, the maximum daily 8-hour average of ozone ( $\text{MDA8-O}_3$ ) was 110.3  
233  $\pm 10.1$  ppb (i.e.,  $236.5 \pm 21.7 \mu\text{g m}^{-3}$ ). The concentration of  $\text{O}_x$  [ $\text{O}_x$  = nitrogen dioxide ( $\text{NO}_2$ ) +  $\text{O}_3$ ] was  
234  $88.6 \pm 29.4$  ppb (Fig. 1B), reflecting a high atmospheric oxidation capacity that drives secondary  
235 transformations of gaseous pollutants (Dou et al., 2024; Xiao et al., 2022). The average  $\text{PM}_{2.5}$  was  $74.3$   
236  $\pm 18.3 \mu\text{g m}^{-3}$  (Fig. 1C). As shown in Fig. 1D – 1E, the average mass concentration of secondary  
237 inorganic aerosol (SIA) in  $\text{NR-PM}_1$  was  $21.3 \pm 4.8 \mu\text{g m}^{-3}$ , with sulfate and nitrate contributing almost  
238 equally to particle mass (i.e., 18.1% and 17.3% respectively). Organic matter in  $\text{NR-PM}_1$  had an



average mass fraction of 56.4% (Fig. 1E). The mass concentrations of primary organic aerosol (POA) and secondary organic aerosol (SOA) were estimated based on the positive matrix factorization (PMF) analysis (Ulbrich et al., 2009). As shown in Fig. 1F, SOA dominated organic matter, contributing an average of 68.9%. Overall, this pollution episode was led by secondary oxidation processes and featured by high contributions of secondary particulate species.



**Figure 1:** Time series of (A) wind direction, wind speed, and relative humidity (RH), (B) mass concentrations of ozone (O<sub>3</sub>) and O<sub>x</sub> (O<sub>3</sub> + nitrogen dioxide, NO<sub>2</sub>), (C) mass concentrations of fine particles (PM<sub>2.5</sub>) and non-refractory submicron particles (NR-PM<sub>1</sub>), and (D) mass fractions of chemical composition of NR-PM<sub>1</sub> are shown. The gray area represents the pollution episode lasting from September 29<sup>th</sup> to October 3<sup>rd</sup>. The red line indicates the sampling time for the individual particle sample. (E) Pie chart showing the average mass fractions of chemical composition of NR-PM<sub>1</sub> during the pollution period. The number in the first row of each part is the average mass concentration and standard deviation (SD) with a unit of μg m<sup>-3</sup>. The number in the second row is the average mass fraction. (F) Pie chart showing the average mass contributions of primary organic aerosol (POA, light green) and secondary organic aerosol (SOA, dark green) to the total organic. Average mass fraction and SD are marked in the pie chart.

### 3.2 Chemical maps of individual particles

Chemical maps of individual particles under dry conditions are displayed in Fig. 2. Different images denote particles located in different regions of interest (ROI, 12 in total) of the sampling substrate. 197 individual particles were investigated in total. Detailed images of total carbon, inorganic, and doubly bonded carbon maps are available in the SI (Fig. S3 – S5). Most submicron particles on the

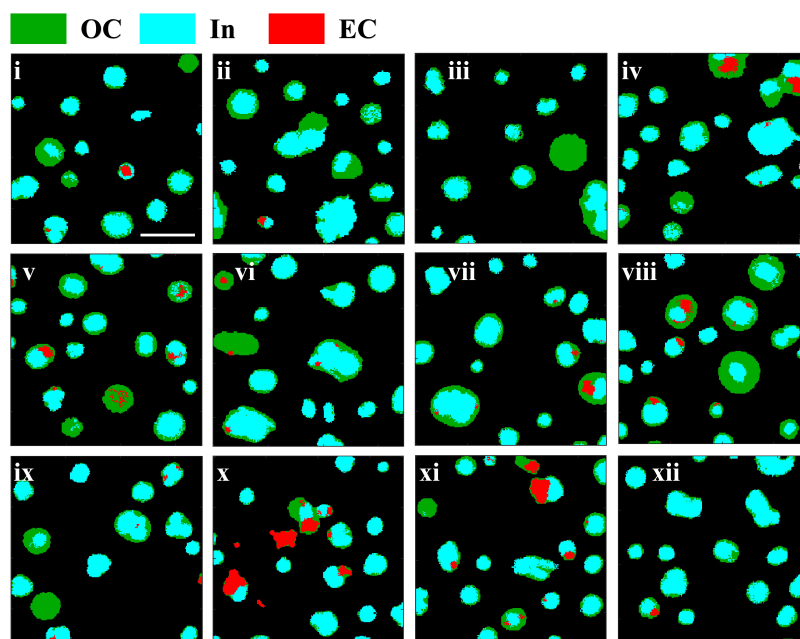


260 substrate were round or nearly round, while supermicron particles predominantly exhibited irregular  
261 shapes. The circular equivalent diameter of individual particles was calculated, with the methods  
262 detailed in the SI. The normalized size distribution of overall particles followed a normal distribution,  
263 with a mean diameter  $\pm$  standard deviation (SD) being  $0.83 \pm 0.30 \mu\text{m}$  (Fig. S6A). A significant  
264 proportion of the particles were within the  $0.4 - 1.2 \mu\text{m}$  size range.

265 As displayed in Fig. 2, chemical maps of individual particles showed that they were dominated by  
266 inorganic substances (colored in cyan), which were likely sulfate that was frequently observed by AMS  
267 (Fig. 1C). Approximately one quarter (24.9%) of the particles contained EC/soot (colored in red),  
268 found either near the center or at the edge of individual particles. Notably, around 82% of these  
269 soot-containing particles had soot located at particle edges. The possible reason is that inorganic  
270 species (such as crystals) pushed soot away from the center of the particles during their efflorescence  
271 (Moffet et al., 2016). The morphology of soot varied, showing fractal or compact structures of various  
272 sizes. Additionally, several particles contained multiple soot components, which was also observed  
273 before (Moffet et al., 2016).

274 Typically, inorganic components and/or soot were encased in organic matter, forming a core-shell  
275 structure characterized by an inorganic-dominated core and an organic-dominated shell. Figure 2  
276 illustrates that most organic-inorganic internally mixed particles exhibited thin coatings, likely from  
277 fresh emissions. Conversely, a few particles have thick coatings, which is indicative of aging processes  
278 in a highly active photochemical environment. Previous studies suggest that most of the  
279 soot-containing particles with thin coatings would have rather smaller absorption enhancement  
280 compared with those with thick coatings (Bond et al., 2006; Moffet et al., 2016).

281 The observed core-shell morphology could also result from liquid-liquid phase separation (LLPS),  
282 influenced by fluctuating ambient RH (Fig. 1A) and determined by the oxygen-to-carbon (O:C) ratio of  
283 the organic fraction (Freedman, 2020; Li et al., 2021; You et al., 2012; You et al., 2014; Freedman,  
284 2017). To test this hypothesis, the O:C ratio of individual particles composed of pure organic  
285 composition was estimated as  $0.53 \pm 0.15$  based on the STXM data. The estimation methods were  
286 displayed in the SI. This falls within the threshold range for LLPS occurrence in ammonium sulfate -  
287 organic mixing particles ( $0 < \text{O:C} < 0.57$ ) (You et al., 2013).



**Figure 2: Chemical maps of individual particles in 12 regions of interest (ROI) of sampling substrate under dry conditions on the basis of pixels. Green, cyan, and red color represent dominant components of organic carbon (OC), inorganic matter (In), and elemental carbon (EC), respectively. The scale bar in the upper left image represents 2  $\mu\text{m}$  and applies to all images.**

Statistically, the particles were categorized into four types based on their mixing state, including pure organic (OC), organic internally mixed with soot (OCEC), organic internally mixed with inorganic (OCIn), and organic internally mixed with inorganic and soot (OCInEC). OCIn particles were the most abundant type in the examined particle population ( $73.1 \pm 7.4\%$ ), followed by OCInEC ( $20.8 \pm 6.7\%$ ) and OCEC ( $4.1 \pm 3.3\%$ ), indicating a highly internally mixed particle population. Pure organic particles only accounted for  $2.0 \pm 2.3\%$ . The calculation of the margin of error of the mixing state proportions can be found in the SI. The mean diameters of OCEC, OCIn, and OCInEC particles were 0.66, 0.79, and 1.02  $\mu\text{m}$ , respectively (Fig. S7). This suggests that the internally mixed particles containing three species families tend to be larger than those composed of two species families.

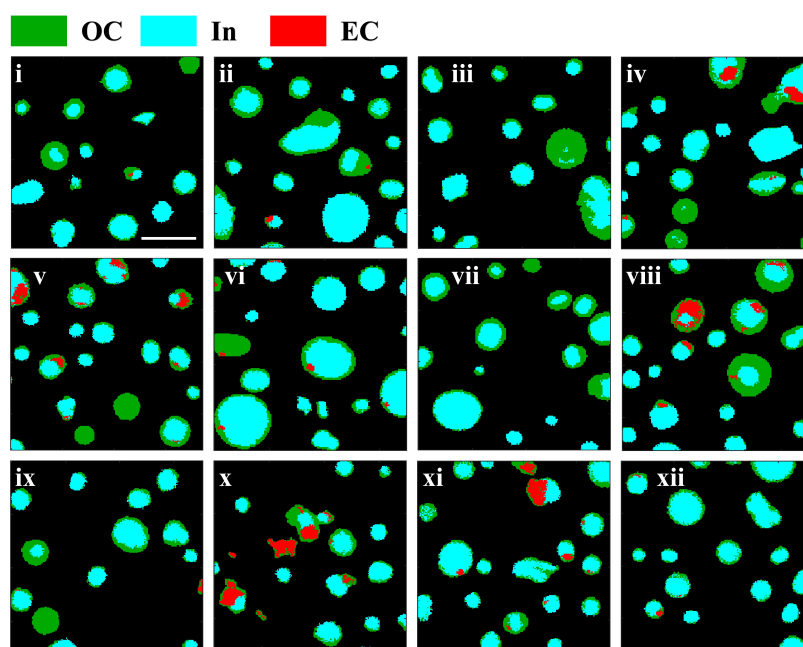
### 3.3 The effects of particle water uptake on chemical maps

Chemical maps of individual particles under humid conditions ( $\text{RH} = 86\%$ ) measured in the environmental cell were displayed in Fig. 3, and these ROI are identical and matched one by one to



those in Fig. 2. It was observed that many particles tended to be more rounded due to water uptake at high RH, especially for particles with diameters in the supermicron range (e.g., particles in (ii), (vi), (xi), and (xii) in Fig. 3). If particles take up significant amounts of water and are homogeneously mixed, they would appear as dominated by inorganic (colored in cyan) due to absorption of a large amount of water at the carbon pre-edge. In contrast, most particles remained inhomogeneous and exhibited a core-shell structure under humid conditions. A possible reason is that the settled RH may not reach the mixing relative humidity (MRH) of particles, which is defined as a threshold where different phases in an aqueous particle mix into one homogeneous phase. This MRH usually varies from 84% to over 90% (Li et al., 2021; You et al., 2014; Zhang et al., 2022).

Additionally, around 87% of soot was located at the edge of the humidified particles, with no obvious location change of soot observed in most particles. This phenomenon aligns with a previous study which indicates that the phase transition of phase-separated particles without phase mixing will not cause the redistribution of soot within individual particles (Zhang et al., 2022).



**Figure 3: Chemical maps of individual particles in 12 ROI of sampling substrate under humid conditions (RH = 86%) measured in an in-situ environmental cell. Green, cyan, and red color represent dominant components of OC, In, and EC, respectively. The scale bar in the upper left image represents 2  $\mu$ m and applied to all images.**



323 Comparing size distributions of particle populations under dry (Fig. S6A) and humid conditions  
324 (Fig. S6B) reveals that they exhibited similar distribution characteristics. The mean diameter of overall  
325 particles at high RH was  $0.86 \pm 0.33 \mu\text{m}$ , compared with  $0.83 \pm 0.30 \mu\text{m}$  under dry conditions. This  
326 indicates that the overall size distribution of the humidified particles shifted a little towards larger  
327 particles due to water uptake. Specifically, approximately 56.3% of the particles showed an average  
328 increase of 14.9% in diameter, while the remaining exhibited an average decrease of 8.2%. Pöhlker et  
329 al. (2014) also observed this abnormal phenomenon where some particles decreased in size with  
330 increasing RH. They suggested that it could be attributed to the decreasing viscosity and increasing  
331 surface tension due to particle water uptake at high RH. This led to larger contact angles between the  
332 collected particles and the substrate, causing the particles to 'bead up' and therefore reducing their  
333 cross-section areas in the view (Pöhlker et al., 2014). In addition, one should note that a small number  
334 of particles at the edge of the ROI did not entirely enter the field of view due to the limited observation  
335 range, which may slightly affect the quantification of their size.

336 According to the criterion for water uptake by individual particles based on the total oxygen  
337 absorbance described in the Sect. 2.3, 41.6% of the particles took up water. As shown in Fig. S8A,  
338 OCIn particles were the dominant mixing state type taking up water (76.8%), followed by OCInEC  
339 (17.1%). There were also several OCEC (2.4%) and OC (3.7%) particles displayed water uptake.  
340 Different particle mixing state types exhibited distinct patterns of hygroscopic behavior. For instance,  
341 43.8% of OCIn particles took up water, while 34.1% of OCInEC particles performed the same. This  
342 difference may be attributed to the varying hygroscopicity of different components. For example, the  
343 single hygroscopic parameter ( $\kappa$ ) of ammonium nitrate, ammonium sulfate, ammonium hydrogen  
344 sulfate, POA, and SOA is 0.58, 0.48, 0.56, 0, and 0.1, respectively (Wu et al., 2016). Based on the  
345 AMS data,  $\kappa$  of bulk aerosols during the sampling period ( $0.25 \pm 0.01$ ) was calculated according to the  
346 Zdanovskii-Stokes-Robinson (ZSR) mixing rule (Stokes and Robinson, 1966), indicating a relatively  
347 low hygroscopic capacity of NR-PM<sub>1</sub> during sampling, which could explain why only less than half of  
348 the particles exhibited water uptake at such high humidity conditions. In addition, the average diameter  
349 of particles taking up water increased from  $0.82 \pm 0.33 \mu\text{m}$  to  $0.91 \pm 0.36 \mu\text{m}$ . The relative frequency  
350 distribution and the size-resolved fraction of particles taking up water can be found in Fig. S8B.



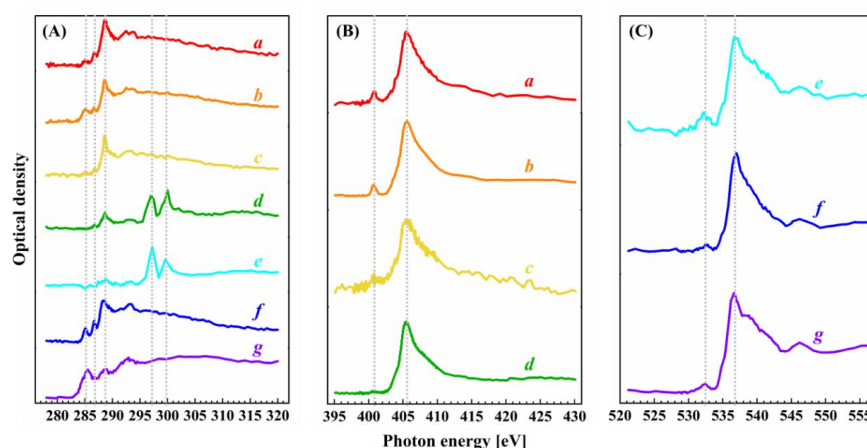
### 3.4 Chemical composition of ambient submicron particles

NEXAFS spectra with high energy resolution were measured at the C (278 – 320 eV), N (395 – 430 eV), and O (525 – 550 eV) K-edges. As shown in Fig. 4A, three notable absorption peaks at the C K-edge were observed at 285.4, 286.7, and 288.6 eV. According to previous literatures (Warwick et al., 1998; Moffet et al., 2010a), the peak at 285.4 eV refers to the characteristic transition of  $sp^2$  hybridized carbon ( $C\ 1s \rightarrow \pi^*_{R(C^*=C)R}$ ). The peak at 286.7 eV may result from the transition of ketonic carbonyl ( $C\ 1s \rightarrow \pi^*_{R(C^*=O)R}$ ), representing ketone and ketone-like compounds. The peak appearing at 288.6 eV represents the characteristic transition of carboxylic carbonyl functional groups ( $C\ 1s \rightarrow \pi^*_{R(C^*=O)OH}$ ), which refers to organic matter and is always found in the outer shell of particles (Moffet et al., 2016; Prather et al., 2013). In addition, two other peaks at 296.8 and 299.6 eV were present (see spectra (d) and (e) in Fig. 4A), corresponding to the  $L_2$ - and  $L_3$ -edges of potassium (K) (Moffet et al., 2010a). In our sample, K may come from biomass burning processes based on a previous study (Wu et al., 2017).

Nitrogen K-edge spectra in Fig. 4B illustrate that ammonium salts were the main nitrogen species in the sample. We observed a broad main peak centered at 405.7 eV, which is the feature of ammonium (Ekimova et al., 2017). A smaller peak was observed at 401.0 eV, which is absorption due to nitrogen gas ( $N_2$ ) either trapped in the inorganic crystal or formed under X-ray exposure (Latham et al., 2017). Absorption of nitrate ( $NO_3^-$ ) and nitrite ( $NO_2^-$ ) commonly have narrow peaks at 405.1 eV and 401.7 eV (Smith et al., 2015), respectively, which were not apparent in our spectra. This is likely because particulate nitrite is below the detection limit, or its peak is masked by the pronounced absorption of ammonium ( $NH_4^+$ ). The solid ammonium nitrate and sodium nitrate salts could exhibit a peak at around 415.0 eV (Smith et al., 2015). However, this was not observed in Fig. 4B. Organic compounds containing nitrogen, such as amino acids, N-heterocyclics, and nitroaromatic compounds, can be abundant in urban aerosol particles due to combustion sources (Yu et al., 2024). They have a large variety of possible peak positions, heights, and widths (Leinweber et al., 2007), making the identification of these compounds difficult. Although a positive identification of specific organic nitriles cannot be made, we note that amino acids and 5- or 6-ring heterocycles commonly have narrow peaks at around 401 eV and broad peaks at 405 eV (Leinweber et al., 2007). We expect that organic nitriles did contribute to the observed N K-edge spectra, although a targeted study on molecular identification would be necessary to establish further certainty.



Oxygen K-edge spectra in Fig. 4C exhibited a large peak at 536.9 eV, which is a representative characteristic of sulfate-rich particles, consistent with the result of AMS. A smaller peak was observed at 532.5 eV, confirming the presence of ketone, aldehyde, or carboxyl functionalities (Colberg et al., 2004; Slowik et al., 2011; Mikhailov et al., 2015; Pöhlker et al., 2014), which aligns with the results from C K-edge spectra. These compositions tend to take up water under humid conditions.



**Figure 4: NEXAFS spectra for individual particles at (A) carbon (C), (B) nitrogen (N), and (C) oxygen (O) K-edges.** In panel (A), peaks were observed at 285.4, 286.7, and 288.6 eV, and two typical peaks appeared at around 296.8 and 299.6 eV in spectra (d) and (e). In panel (B), a main peak appeared at 405.7 eV, and a smaller peak appeared at 410.0 eV. In panel (C), a main peak appeared at 536.9 eV, and a smaller peak was at 532.5 eV. Each small case letter of a spectrum stands for the average result of all the pixels within an individual particle. The same letter in different panels doesn't refer to a same particle.

#### 4 Conclusions and implications

Particles in the atmosphere usually act as reaction vessels for heterogeneous reactive uptake of gaseous molecules, which play an important part in gas-particle partitioning and secondary aerosol formation (Abbatt et al., 2012; Davidovits et al., 2011; Kolb et al., 2010). However, determining the particle physicochemical properties is crucial but challenging due to the complexity and inhomogeneity of aerosols particles (Barbaray et al., 1979; Zong et al., 2022). So far, there is a lack of study on direct observation of the physicochemical properties of urban aerosols at the single particle level under different conditions, which hinders our understanding of the role of urban particle aerosols in multiphase and heterogeneous chemistry.

In this study, we used STXM/NEXAFS spectroscopy combined with an environmental cell to





402 image and quantify the chemical morphology and water uptake behavior of individual submicron  
403 particles collected in an urban pollution atmosphere. Results show that most organic compounds were  
404 internally mixed with inorganic and/or soot, generally presenting a core-shell structure with an  
405 inorganic core and an organic shell. Internally mixed particles composed of organic carbon and  
406 inorganic matter dominated the particle population by  $73.1 \pm 7.4\%$ . At 86%RH, 41.6% of the particles  
407 took up water, with OCIn particles making up 76.8% of these hygroscopic particles. The relatively low  
408 hygroscopicity of bulk aerosols during the sampling period ( $\kappa = 0.25 \pm 0.01$ ) helps to explain the  
409 reason why only less than half of the particles took up water. Besides, the majority of particles still  
410 showed a heterogeneous core-shell morphology under humid conditions.

411 This study directly displays the dominant chemical morphology (i.e., core-shell structure) and  
412 hygroscopic behavior of individual submicron urban aerosol particles at the microscale. Results  
413 highlight the importance of taking the core-shell structure into consideration for estimating the uptake  
414 coefficient and investigating heterogeneous reactions, which can improve our comprehension of  
415 atmospheric processes of secondary aerosols in typical urban pollution atmospheres. Moreover,  
416 previous studies found that the impact of organics on the reactive uptake coefficient mainly relies on  
417 organic species, organic content, and particle mixing state (Wagner et al., 2013). Therefore, it is  
418 instrumental to conduct research on the basis of organic molecules and functional groups in further  
419 study.



420 **ASSOCIATED CONTENT**

421 **Data availability.** The data presented in this article can be accessed through the corresponding author  
422 Zhijun Wu via E-mail ([zhijunwu@pku.edu.cn](mailto:zhijunwu@pku.edu.cn)).

423 **Author contributions.** YSZ, PAA, BBW, and JD measured the individual particle sample by  
424 STXM/NEXAFS. YSZ, ZJW, YZ, YLG, QC, and SYC carried out the field observation and obtained  
425 data. RQM and PAA processed and analyzed data. All authors discussed the results and contributed to  
426 the writing of this paper. RQM prepared the manuscript. ZJW, PAA, JC, XRK, MA, and MH further  
427 modified and improved the manuscript.

428 **Competing Interests.** The authors declare that they have no conflict of interest.

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