

1 **Atmospheric chemical processing dictates aerosol aluminum solubility: insights from**
2 **field measurement at two locations in Northern China**

3

4 Tianyu Zhang,^{1,5} Yizhu Chen,^{1,5} Huanhuan Zhang,² Lei Liu,³ Chengpeng Huang,⁴ Zhengyang
5 Fang,^{1,a} Yifan Zhang,^{1,5} Fu Wang,⁴ Lan Luo,⁴ Guohua Zhang,¹ Xinming Wang,¹ Mingjin
6 Tang^{1,6,*}

7

8 ¹ State Key Laboratory of Advanced Environmental Technology and Guangdong Key
9 Laboratory of Environmental Protection and Resources Utilization, Guangzhou Institute
10 of Geochemistry, Chinese Academy of Sciences, Guangzhou, China

11 ² Guangzhou Marine Geological Survey, China Geological Survey, Guangzhou, China

12 ³ Hangzhou International Innovation Institute, Beihang University, Hangzhou, China

13 ⁴ Longhua Center for Disease Control and Prevention of Shenzhen, Shenzhen, China

14 ⁵ College of Earth and Planetary Sciences, University of Chinese Academy of Sciences, Beijing,
15 China

16 ⁶ Institute of Surface-Earth System Science, School of Earth System Science, Tianjin
17 University, Tianjin, China

18 ^a Current address: Institute of Low Temperature Science, Hokkaido University, Sapporo, 060-
19 0819, Japan

20 Correspondence: Mingjin Tang (mingjintang@126.com)

21

22

23 **Abstract**

24 Deposition of mineral dust aerosol into open oceans impacts marine biogeochemistry, and
25 the deposition flux can be constrained using dissolved aluminum (Al) in surface seawater as a
26 tracer. However, aerosol Al solubility, a critical parameter used in this method, remains highly
27 uncertain. We investigated seasonal variations of aerosol Al solubility for supermicron and
28 submicron particles at two locations (Xi'an and Qingdao) in Northern China. Aerosol Al
29 solubility was very low at Xi'an, showed no apparent variation with seasons or relative
30 humidity, and was not correlated with sulfate or nitrate; in contrast, Al solubility was much
31 higher at Qingdao, exhibited distinct seasonal variability, and increased with relative humidity
32 and the abundance of sulfate and nitrate. All these features observed for Al solubility at the two
33 locations can be explained by the effects of atmospheric chemical processing. Mineral dust
34 transported to Xi'an (an inland city in Northwest China) was still not obviously aged and thus
35 chemical processing had little effect on aerosol Al solubility; after arriving at Qingdao (a
36 coastal city in the Northwest Pacific), mineral dust was substantially aged by chemical
37 processing, leading to significant enhancement in aerosol Al solubility. Our work further
38 reveals that aerosol liquid water and acidity play vital roles in the dissolution of aerosol Al by
39 atmospheric chemical processing. We suggest that chemical aging can lead to spatiotemporal
40 variation of aerosol Al solubility, and this should be considered when using dissolved Al in
41 surface seawater to constrain oceanic dust deposition.

42

43

44 **1. Introduction**

45 As an important type of tropospheric aerosols, mineral dust aerosol greatly impacts
46 atmosphere chemistry, climate, and ecological systems (Jickells et al., 2005; Tang et al., 2016;
47 Kok et al., 2023). After long-range transport, deposition of mineral dust into the oceans is a
48 major external source of several nutrient and toxic elements for surface seawater (Moore et al.,
49 2013; Westberry et al., 2023), impacting primary production and biogeochemical cycles in the
50 oceans and having further feedback on the climate system (Mahowald, 2011; Jiang et al., 2024).
51 The deposition flux of mineral dust aerosol into the oceans should be accurately estimated
52 before we can assess its impacts on marine biogeochemistry in a reliable manner (Schulz et al.,
53 2012; Anderson et al., 2016). Previous studies used several different methods to estimate dust
54 deposition fluxes and found large discrepancies (Huneeus et al., 2011; Anderson et al., 2016).

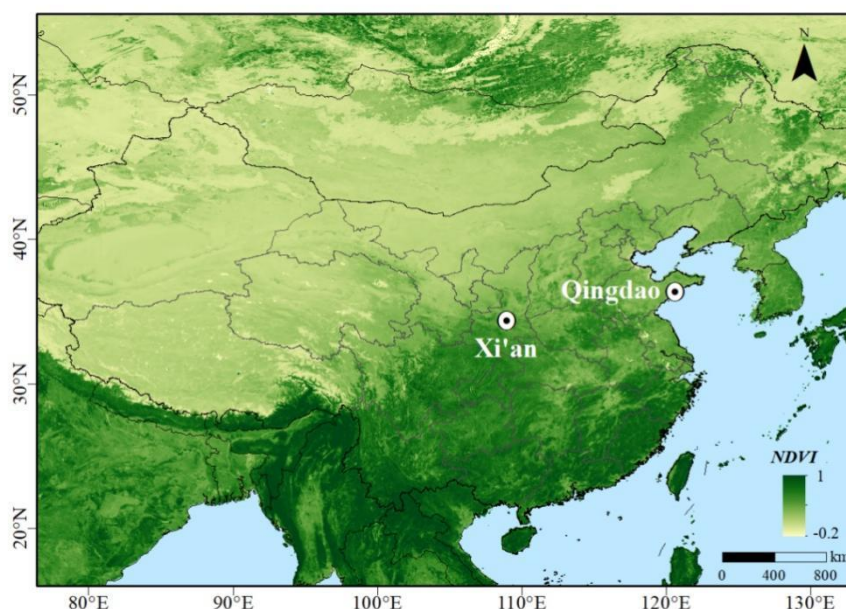
55 Deposition of mineral dust aerosol is the dominant source of dissolved aluminum (Al) in
56 the surface water of open oceans, and dissolved Al is generally considered to be chemically
57 and biologically inactive in seawater. As a result, dissolved Al concentrations in surface
58 seawater could be used to calculate dust deposition flux into the oceans (Measures and Brown,
59 1996; Measures and Vink, 2000), and the fractional solubility of aerosol Al (the fraction of
60 aerosol Al that can be dissolved) is one of the key parameters used in this method. Previous
61 studies which used this method to estimate dust depositions fluxes (Han et al., 2008; Measures
62 et al., 2010; Grand et al., 2015; Benaltabet et al., 2022) usually assumed uniform Al solubility
63 values in the range of 1.5-5%. However, field measurements found that aerosol Al solubility
64 could vary by more than an order of magnitude (Baker et al., 2006; Buck et al., 2013), and
65 thereby using a uniform aerosol Al solubility value could lead to large uncertainties in

66 estimated dust deposition fluxes (Han et al., 2008; Xu and Weber, 2021). In order to better
67 constrain the oceanic dust deposition using dissolved Al in seawater as a tracer, we need to
68 develop parameterizations for aerosol Al solubility, and this requires spatiotemporal variability
69 of aerosol Al solubility to be understood and processes and mechanisms which drive such
70 variations to be elucidated.

71 The initial Al solubility is generally low (typically <1.5%) for soil or mineral dust samples
72 (Mulder et al., 1989; Duvall et al., 2008; Shi et al., 2011; Aghnatios et al., 2014; Li et al., 2022),
73 and field studies found that aerosol Al solubility in the troposphere could be much higher and
74 showed wide variability. For example, Al solubility ranged from 0.2-15.9% for total suspended
75 particles (TSP) over the Pacific (Buck et al., 2013), and were in the range of 3-78% over the
76 Atlantic (Buck et al., 2010; Chance et al., 2015). Some studies (Measures et al., 2010; Sakata
77 et al., 2023) found good correlations between dissolved aerosol Al (or Al solubility) and acid
78 species in aerosol particles, and thus suggested that chemical processes in the atmosphere could
79 substantially enhance aerosol Al solubility; furthermore, Li et al. (2017) found that Al solubility
80 was remarkably increased during cloud events when cloud processing enhanced the formation
81 of secondary inorganic ions (mainly sulfate and nitrate) and thus increased the acidity of cloud
82 droplets. However, Yang et al. (2023) found no correlations between Al solubility and the
83 concentrations of aerosol acidic species, and concluded that the effect of acid processing on Al
84 solubility was negligible. Aerosol Al solubility over the Atlantic appeared to be higher for air
85 masses from Europe than those from the Saharan region (Baker et al., 2006; López-García et
86 al., 2017), and some studies hypothesized that this could be potentially explained by the

87 influence of anthropogenic aerosol Al if it had higher solubility than mineral dust (Paris et al.,
88 2010; López-García et al., 2017).

89 It can be concluded that although aerosol Al solubility in the atmosphere was explored by
90 several previous studies, our understanding is still very limited. For example, it remains unclear
91 why aerosol Al solubility shows large spatial and temporal variation. Some work suggested
92 that atmospheric chemical aging could enhance aerosol Al solubility, but the mechanisms and
93 key environmental factors have not been elucidated. Furthermore, the effects of particle size
94 on aerosol Al solubility have not been well understood.



95
96 **Figure 1.** A map of East Asia and surrounding areas. The two locations (Xi'an and Qingdao)
97 where we collected aerosol particles are highlighted. NDVI: normalized difference vegetation
98 index provided by MODIS (Moderate Resolution Imaging Spectroradiometer).

99
100 In this work, we collected supermicron ($>1 \mu\text{m}$) and submicron ($<1 \mu\text{m}$) aerosol particles
101 at Xi'an and Qingdao, both located in Northern China, and investigated seasonal variations of

102 aerosol Al solubility at these two locations. Taklimakan and Gobi Deserts in Northwestern
103 China are two important source regions of Asian dust (Prospero et al., 2002). As shown in
104 Figure 1, Xi'an is an inland city in Northwestern China, located at the southern edge of the
105 Loess Plateau which is also an active source of mineral dust (Cao et al., 2008; Jeong, 2020;
106 Haugvaldstad et al., 2024), and the aging extent of mineral dust at Xi'an was found to be quite
107 limited (Wang et al., 2014; Wu et al., 2017). As Asian dust is transported eastward, it passes
108 over the North China Plain where anthropogenic emission is very high, and may become much
109 more aged when arriving at Qingdao, a coastal city of the Northwest Pacific (Li et al., 2014;
110 Pan et al., 2017). By comparing aerosol Al solubility at Xi'an and Qingdao, our work can
111 provide valuable insights into how and to which extent aging processes during long-range
112 transport can change aerosol Al solubility. Dust aerosol concentrations and meteorological
113 conditions vary remarkably at different seasons in Northern China; as a result, examining its
114 seasonal variations provides a good opportunity to understand the factors which regulate
115 aerosol Al solubility.

116 **2. Materials and methods**

117 **2.1 Sample collection**

118 Samples were collected at two cities (Xi'an and Qingdao) in Northern China at four
119 different seasons during 2021-2023 (Zhang et al., 2023; Chen et al., 2024), and further details
120 can be found in the supplement (Text S1 and Table S1). In brief, supermicron ($>1\ \mu\text{m}$) and
121 submicron ($<1\ \mu\text{m}$) particles were simultaneously collected using a two-stage aerosol sampler
122 (TH-150C, Tianhong Co., China) which was operated at 100 L/min, and the sampling duration
123 was typically 23.5 hours for each pair of aerosol samples. Whatman 41 cellulose filters were

124 used for aerosol collection in our work, and they were acid-washed before being used for
125 aerosol sampling to reduce background levels (Zhang et al., 2022). A total of 126 and 106 pairs
126 of aerosol samples were collected at Xi'an and Qingdao, respectively (Zhang et al., 2023; Chen
127 et al., 2024). After collection, all the aerosol samples were stored at -20°C for further analysis.

128 In addition to aerosol particles, we also sampled atmospheric acidic and alkaline gases
129 (mainly NH_3 , HCl and HNO_3) at Qingdao, using a ChemComb 3500 Speciation Collection
130 Cartridge (Thermo Fisher Scientific, USA) at a flow rate of 10 L/min (Walters and Hastings,
131 2018; Fang et al., 2025). Gas sampling was carried out concurrently with aerosol sampling. In
132 brief, NH_3 , HNO_3 and HCl were absorbed onto the inner walls of two tandem honeycomb
133 diffusion tubes coated with proper adsorbents, and then converted into NH_4^+ , NO_3^- and Cl^- .
134 After the sampling was completed, 20 mL ultrapure water was used to rinse each tube
135 immediately, and a PTFE membrane syringe filter ($0.22\ \mu\text{m}$ in pore size) was used to filter the
136 solution. The solution was then frozen at -20°C for further analysis.

137 **2.2 Sample analysis and aerosol acidity calculation**

138 Sample pretreatment and analysis were detailed in our previous work (Zhang et al., 2022),
139 and therefore are only briefly summarized here. The first half of a filter (and only one quarter
140 of a filter for supermicron particles) was shredded and digested in a Teflon jar using a
141 microwave digestion instrument. After digestion, the Teflon jar was filled with 1% HNO_3 (20
142 mL), and a PTFE membrane syringe filter ($0.22\ \mu\text{m}$ in pore size) was used to filter the solution;
143 subsequently, the solution was analyzed by inductively coupled plasma-mass spectrometry
144 (ICP-MS) to determine total concentrations of individual trace elements, including Al.

145 The other half of a filter was immersed in ultrapure water (20 mL) and stirred using an

146 orbital shaking for two hours; in the next step, the solution was filtered using a PTFE membrane
147 syringe filter (0.22 μm in pore size) and divided into two parts. The first solution was acidified
148 to contain 1% HNO_3 and subsequently analyzed by ICP-MS to determine the concentrations
149 of dissolved trace elements; the second solution was analyzed by ion chromatography (IC) to
150 quantify the concentration of water-soluble cations and anions.

151 The solutions obtained from honeycomb diffusion tubes (see Section 2.1 for more details)
152 were also analyzed using IC to determine the concentrations of gaseous NH_3 , HCl and HNO_3
153 in the atmosphere. ISORROPIA-II, a widely used aerosol thermodynamic model (Fountoukis
154 and Nenes, 2007), was employed in this work to calculate the acidity of supermicron and
155 submicron particles. It was operated in the forward mode, and aerosol particles were assumed
156 to remain metastable. Input parameters included concentrations of water-soluble ions in aerosol
157 particles and gaseous NH_3 , HCl and HNO_3 , temperature and relative humidity (RH). Our
158 previous work found good agreement between measured and calculated NH_3 partitioning
159 coefficients at Qingdao (Fang et al., 2025), and as a result the method we used could well
160 estimate the acidity of supermicron and submicron particles.

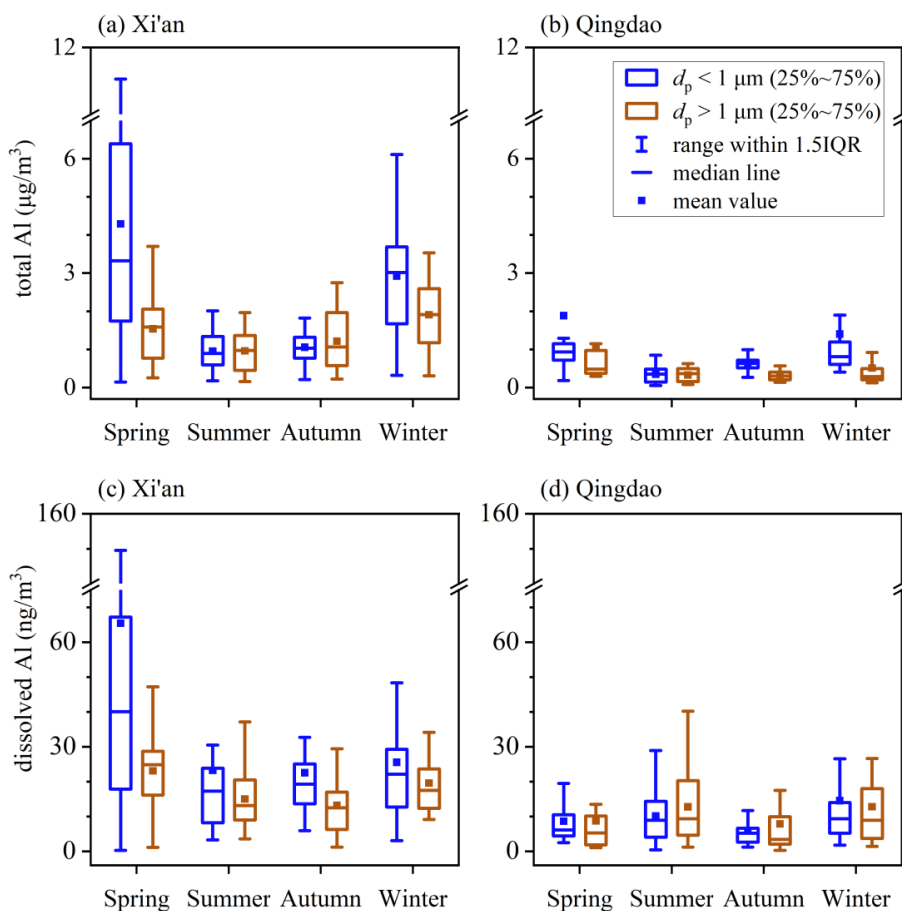
161 **3. Results**

162 **3.1 Seasonal variations of total and dissolved aerosol Al**

163 **3.1.1 Total aerosol Al**

164 Figure 2 displays seasonal variations of total and dissolved aerosol Al at Xi'an and
165 Qingdao. At Xi'an (Figure 2a), total Al in supermicron particles showed highest concentrations
166 in spring and winter (1.54 ± 0.89 and $1.91 \pm 0.93 \mu\text{g}/\text{m}^3$) and lowest concentrations in summer
167 ($0.96 \pm 0.54 \mu\text{g}/\text{m}^3$); a similar seasonal pattern was observed for submicron particles, with total

168 Al concentrations being highest in spring and winter (4.29 ± 3.70 and 2.92 ± 1.47 $\mu\text{g}/\text{m}^3$) and
 169 lowest in summer (0.95 ± 0.44 $\mu\text{g}/\text{m}^3$). At Qingdao (Figure 2b), total Al concentrations in
 170 supermicron particles were highest in spring (1.04 ± 1.12 $\mu\text{g}/\text{m}^3$) and lowest in summer and
 171 autumn (0.33 ± 0.18 and 0.31 ± 0.12 $\mu\text{g}/\text{m}^3$); similarly, for submicron particles, total Al
 172 concentrations were also highest in spring (1.88 ± 2.51 $\mu\text{g}/\text{m}^3$) and lowest in summer and
 173 autumn (0.35 ± 0.22 and 0.65 ± 0.82 $\mu\text{g}/\text{m}^3$). For each season the median concentration of total
 174 aerosol Al was usually higher in submicron particles than supermicron particles at both
 175 locations (and there were some exceptions, as shown in Figures 1a and 1b). This is related to
 176 size dependence of mineralogy and elemental compositions of mineral dust aerosol, which is
 177 not well studied and deserves further investigation.



178
 179 **Figure 2.** Seasonal variations of total and dissolved aerosol Al for submicron and supermicron

180 particles: (a) total Al at Xi'an; (b) total Al at Qingdao; (c) dissolved Al at Xi'an; (d) dissolved
181 Al at Qingdao.

182

183 Overall, total aerosol Al concentrations showed similar seasonal variations at Xi'an and
184 Qingdao, being highest in spring and lowest in summer. This was consistent with previous
185 studies carried out in other locations in East Asia, such as Zhengzhou (Wang et al., 2019),
186 Beijing (Zhang et al., 2013), Huaniao Island in the East China Sea (Guo et al., 2014), and Japan
187 (Sakata et al., 2023). In East Asia, mineral dust aerosol was emitted into the atmosphere mainly
188 in spring, leading to the increase in total aerosol Al concentrations. Lowest concentrations of
189 total aerosol Al were observed in summer because precipitation in Northern China mainly
190 occurred in summer, leading to enhanced wet deposition of aerosol particles (Cao and Cui,
191 2021). Furthermore, Qingdao was frequently affected by marine air masses in summer, and
192 this is also one reason why total aerosol Al concentrations were lower in summer than other
193 seasons. Total aerosol Al concentrations were higher in winter than summer and autumn at
194 Xi'an, and one major reason is that meteorological conditions favored the accumulation of
195 aerosol particles (including aerosol Al) during winter (Cao and Cui, 2021). **Futhermore, besides**
196 **spring, Asian dust also occurs in winter (Cai et al., 2020; Wang et al., 2020), and a previous**
197 **study (Huang et al., 2014) suggested that the dust-related source, including local resuspended**
198 **dust, contributed 56% to PM_{2.5} during a severe haze event at Xi'an.**

199 As summarized in the supplement (Table S2), total aerosol Al concentrations exhibited
200 evident spatial variations in East Asia. As Asian dust was transported eastward to the North
201 Pacific, a clear decrease in aerosol Al concentrations was observed. Mineral dust was the

202 dominant source for aerosol Al, and therefore concentrations of aerosol Al were found to be
203 very high in desert regions. For example, total Al concentrations in TSP could reach $24 \mu\text{g}/\text{m}^3$
204 over the Taklimakan Desert (Zhang et al., 2003). In our current study, annual average total Al
205 concentrations at Xi'an, an inland city close to the desert, were reported to be 1.42 ± 0.86 and
206 $2.28 \pm 2.35 \mu\text{g}/\text{m}^3$ for supermicron and submicron particles, much lower than that observed over
207 the Taklimakan Desert. Further decrease in total Al concentrations was observed in coastal and
208 oceanic regions. For example, our work found that the annual average total Al concentrations
209 were 0.56 ± 0.75 and $1.08 \pm 1.67 \mu\text{g}/\text{m}^3$ for supermicron and submicron particles at Qingdao,
210 lower than those at Xi'an; total Al concentrations in TSP ranged from 0.17 to $1.72 \mu\text{g}/\text{m}^3$ in
211 Hiroshima (Sakata et al., 2023), and further decreased to 1-56 ng/m^3 in Hawaii in the central
212 Pacific (Measures et al., 2010).

213 **3.1.2 Dissolved aerosol Al**

214 At Xi'an (Figure 2c), for supermicron particles, dissolved aerosol Al concentrations were
215 highest in spring ($23.1 \pm 10.9 \text{ ng}/\text{m}^3$) and lowest in summer and autumn (15.0 ± 8.7 and 13.2 ± 8.6
216 ng/m^3); for submicron particles, dissolved Al concentrations were also highest in spring
217 ($65.4 \pm 79.2 \text{ ng}/\text{m}^3$) and lowest in summer and autumn (23.2 ± 23.4 and $22.6 \pm 20.1 \text{ ng}/\text{m}^3$). Total
218 (Figure 2a) and dissolved aerosol Al (Figure 2c) showed similar seasonal patterns at Xi'an,
219 indicating that dissolved aerosol Al was mainly regulated by total aerosol Al.

220 As shown in Figure 2d, the average dissolved aerosol Al concentrations were 8.8 ± 10.8 ,
221 12.8 ± 11.1 , 7.9 ± 10.5 and $12.8 \pm 12.9 \text{ ng}/\text{m}^3$ for supermicron particles at Qingdao in spring,
222 summer, autumn, and winter, respectively, and 8.7 ± 5.8 , 10.2 ± 8.2 , 6.0 ± 4.8 and $14.5 \pm 15.2 \text{ ng}/\text{m}^3$
223 for submicron particles. Dissolved aerosol Al concentrations were highest in summer and

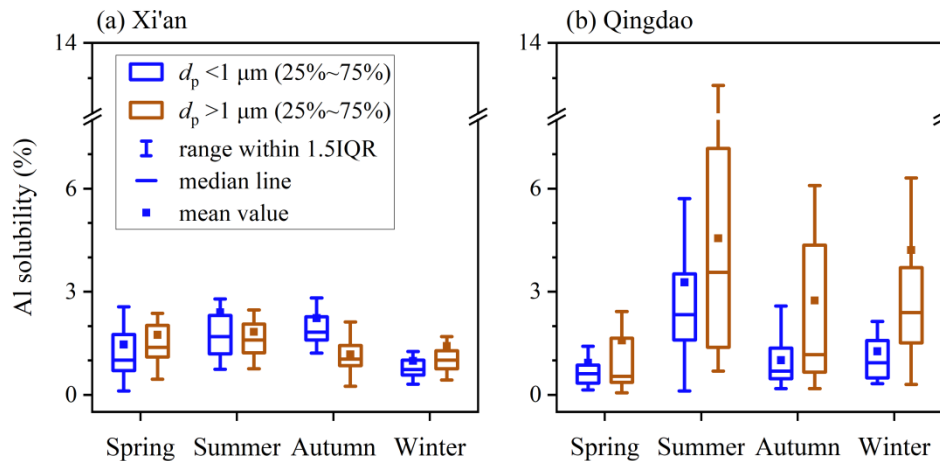
224 winter and lowest in autumn for both supermicron and submicron particles. In contrast to Xi'an,
225 total and dissolved aerosol Al at Qingdao showed different seasonal patterns (Figures 2b and
226 2d); for example, total Al concentrations were lowest in summer at Qingdao when dissolved
227 Al concentrations were highest. This indicates that dissolved aerosol Al at Qingdao was not
228 only regulated by total aerosol Al but also affected by other factors such as atmospheric aging
229 processes.

230 Compared to Xi'an, dissolved Al concentrations at Qingdao were lower across all the four
231 seasons, mainly because total Al concentrations were much lower at Qingdao (Tables S3-S4 in
232 the supplement). As shown in Figure 2, similar seasonal patterns were observed at two
233 locations for total aerosol Al, but dissolved aerosol Al showed very different seasonality; this
234 suggests that seasonal patterns of aerosol Al solubility were different at Xi'an and Qingdao, as
235 presented in Section 3.2.

236 **3.2 Fractional solubility of aerosol Al**

237 **3.2.1 Seasonal variations of Al solubility**

238 Figure 3 displays aerosol Al solubility in different seasons at Xi'an and Qingdao. The
239 median solubilities of aerosol Al were determined to be 1.38%, 1.59%, 1.04% and 1.01% for
240 supermicron particles at Xi'an in spring, summer, autumn and winter, respectively, and 1.01%,
241 1.69%, 1.82% and 0.74% for submicron particles. Aerosol Al solubilities were generally low
242 for the four seasons at Xi'an, showing no apparent variation with seasons (Figure 3a). In
243 contrast, aerosol Al solubilities exhibited distinct seasonal variability at Qingdao (Figure 3b),
244 and the median Al solubilities were highest in summer (3.56% and 2.33%) and lowest in spring
245 (0.54% and 0.61%) for both supermicron and submicron particles.



246

247 **Figure 3.** Seasonal variations of aerosol Al solubility for submicron and supermicron particles

248 at (a) Xi'an and (b) Qingdao.

249

250 In three seasons (summer, autumn and winter), aerosol Al solubility at Qingdao was
 251 higher than that at Xi'an (Figure 3, Table S5). There are several important dust sources in
 252 Northwest China, being far from (up to a few thousand km) or close to Xi'an. More importantly,
 253 anthropogenic emission in Northwest China is much smaller than the North China Plain, and
 254 thus the aging extent of mineral dust transported to Xi'an was rather limited (Wang et al., 2014;
 255 Wu et al., 2017). On the contrary, Qingdao is much farther from deserts; consequently, after
 256 long-distance transport over the North China Plain where anthropogenic emission is very large,
 257 mineral dust aerosol which arrived at Qingdao was substantially aged (Trochkin et al., 2003;
 258 Takahashi et al., 2011; Jeong, 2020), thereby leading to enhanced dissolution of aerosol Al and
 259 thus the increase in Al solubility. Mineral dust from different desert regions and local
 260 suspended dust cannot explain higher Al aerosol solubility observed at Qingdao, as previous
 261 work showed that Al solubility was low for soil samples from different regions (Mulder et al.,
 262 1989; Duvall et al., 2008; Shi et al., 2011; Aghnatiotis et al., 2014; Li et al., 2022).

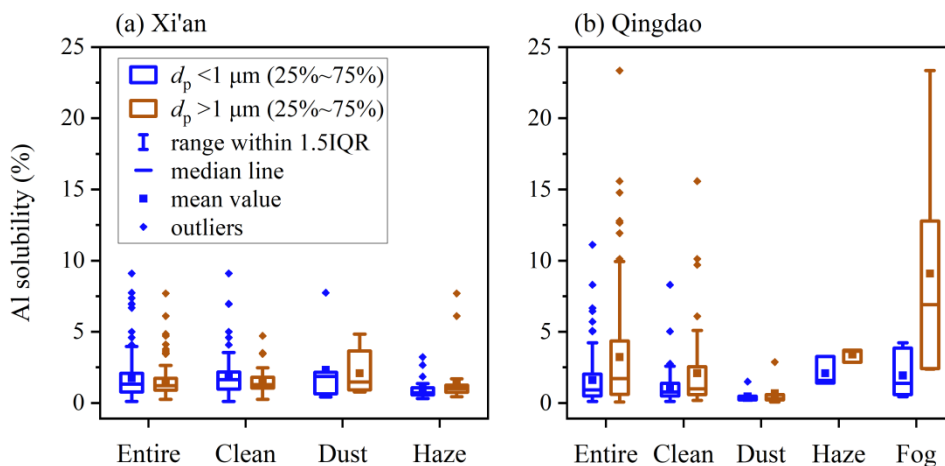
263 On the other hand, no obvious difference in aerosol Al solubility was observed between
264 Xi'an and Qingdao in spring, with median aerosol Al solubilities being <1.4% for supermicron
265 and submicron particles (Figure 3). This agrees with a previous study (Hsu et al., 2010) which
266 found that aerosol Al solubility was very low (average: ~0.7%) in spring even over the East
267 China Sea. Furthermore, similar to what we observed in spring at Xi'an and Qingdao, Al
268 solubility was found to be low (<1.5%) for surface soil particles (Mulder et al., 1989; Duvall
269 et al., 2008; Shi et al., 2011; Aghnatiou et al., 2014; Li et al., 2022). Overall, our work implies
270 that in spring when Asian dust occurred most frequently, mineral dust particles arriving at
271 Qingdao after long-distance transport did not show substantial increase in Al solubility.

272 **3.2.2 Al solubility under different weather conditions**

273 We encountered four representative weather conditions (i.e. clean, dust, haze and fog days)
274 during our sampling at Xi'an and Qingdao, and investigated aerosol Al solubility under
275 different weather conditions (Figure 4, Tables S6-S7).

276 At Xi'an, no apparent difference in Al solubility was observed during clean, haze, and
277 dust days (Figure 4a, Table S6), with median values in the range of 1.01-1.47% for supermicron
278 particles and 0.72-1.86% for submicron particles. Al solubility was found to be <1.2% for three
279 mineral dust samples (Luochuan loess, Arizona test dust, and dust collected during a dust storm
280 in Xinjiang) (Li et al., 2022), and ranged from 0.47% to 1.42% for aerosol particles generated
281 using soil samples from Saharan desert (Shi et al., 2011). Compared to mineral dust in source
282 regions, Al solubility was not higher under different weather conditions at Xi'an. In addition,
283 although emission and accumulation of anthropogenic pollutants was greatly enhanced during
284 haze days at Xi'an (An et al., 2019; Cao and Cui, 2021), there was no obvious increase in

285 aerosol Al solubility, indicating that the effects of anthropogenic emissions on aerosol Al
 286 solubility was limited at Xi'an. Therefore, one may conclude that aerosol Al solubility at Xi'an
 287 was not different from initial Al solubility of mineral dust.



288
 289 **Figure 4.** Aerosol Al solubility under different weather conditions for submicron and
 290 supermicron particles: (a) Xi'an, (b) Qingdao.

291
 292 Being different to Xi'an, aerosol Al solubility at Qingdao shows remarkable variations
 293 under different weather conditions (Figure 4b, Table S7). Median Al solubilities were
 294 determined to be 0.31% and 0.24% for supermicron and submicron particles during dust days,
 295 lower than these on clean days (0.99% and 0.77%, respectively). This is probably because
 296 higher wind speeds during dust events hindered the accumulation of atmospheric pollutants
 297 and shortened the transport time to Qingdao, and thus limiting the aging of mineral dust aerosol.
 298 This explanation is supported by a recent study (Zhang et al., 2024) which found that the aging
 299 extent of dust particles in Japan was much lower during fast-moving dust events than slow-
 300 moving dust events. Moreover, large amounts of alkaline components (such as carbonates)
 301 which were emitted to the atmosphere during dust days neutralized acid species and therefore

302 inhibited acid-promoted dissolution of aerosol trace elements (Zhi et al., 2025). Our work
303 implies that during large dust events increase in aerosol Al solubility may be rather limited
304 when dust is transported to Qingdao; nevertheless, when dust is transported further eastward
305 to the open ocean, atmospheric chemical processing may substantially increase aerosol Al
306 solubility.

307 Figure 4b also suggests that aerosol Al solubilities were much higher during haze and fog
308 days at Qingdao, when compared to clean days. Highest Al solubilities were observed during
309 fog days, with median values being 6.90% for supermicron particles and 1.38% for submicron
310 particles, followed by haze days (3.64% and 1.58%, respectively). This is very likely due to
311 enhanced chemical processing during haze and fog periods (Shi et al., 2020; Shang et al., 2024),
312 and especially during fog days the large increase in RH cause huge increase in aerosol liquid
313 water, therefore greatly promoting aqueous reactions and Al dissolution. Acid and ligand
314 processing can both enhance aerosol Al solubility, although at present it is difficult to
315 disentangle their individual contributions.

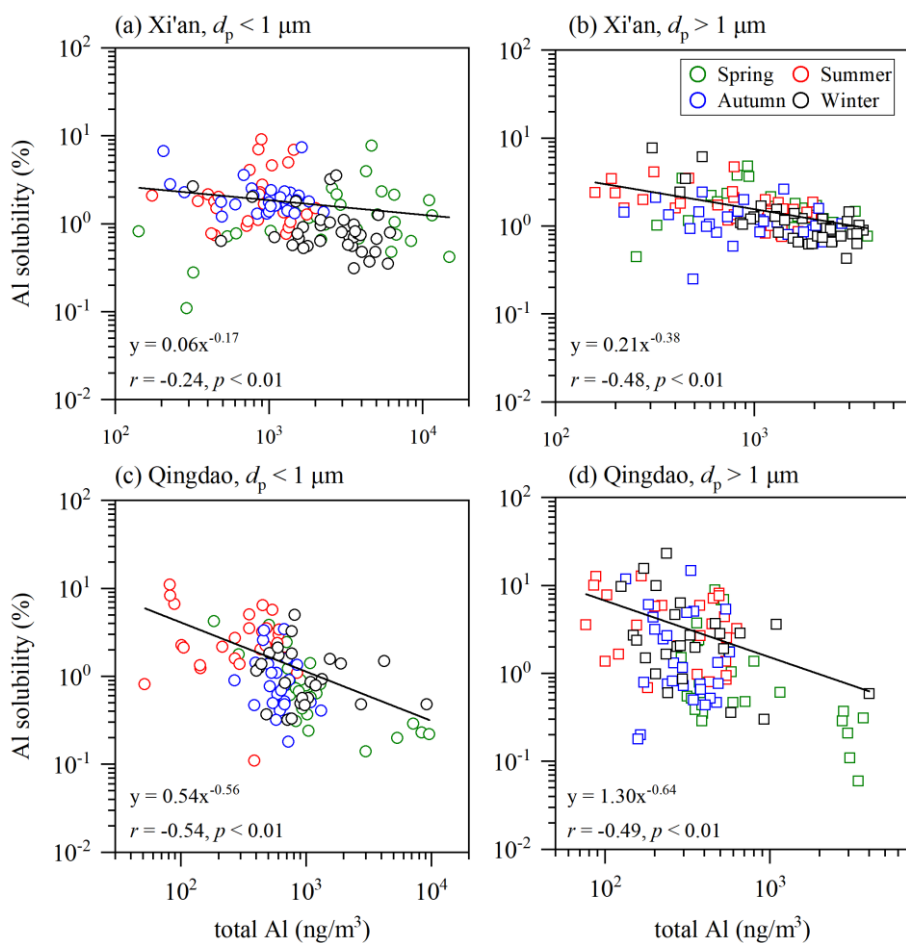
316 In summary, aerosol Al solubility at Xi'an was low in general, and did not show much
317 variability in different seasons or under different weather conditions. Compared to Xi'an,
318 aerosol Al solubility was higher at Qingdao; furthermore, it was higher in the other three
319 seasons than in spring, and much higher for haze and fog days than dust days. These results
320 imply that atmospheric aging had little effect on aerosol Al solubility at Xi'an but could
321 remarkably increase aerosol Al solubility at Qingdao, as further elaborated in Section 4.

322 **4. Discussion**

323 As shown in Figure 5, our work observed the inverse dependence of aerosol Al solubility
324 on total Al concentrations at both Xi'an and Qingdao, given by Eq. (1):

$$325 \quad f_s(\text{Al}) = a \times [\text{Al}]^{-b} \quad (1)$$

326 where $f_s(\text{Al})$ is aerosol Al solubility (%) and $[\text{Al}]$ is total Al concentration (ng/m^3). Such
327 relationship was also reported in some previous studies (Jickells et al., 2016; Shelley et al.,
328 2018; Baker et al., 2020; Shelley et al., 2025). Baker and Jickells (2006) suggested that such
329 inverse relationship was due to that larger particles have higher deposition velocities and lower
330 Al solubility: aerosol Al concentrations decrease during transport in the atmosphere due to
331 deposition, with deposition being faster for larger particles; as a result, aerosol particles will be
332 enriched with smaller particles with higher Al solubility. However, Shi et al. (2011) found no
333 substantial change in Al solubility with particle size for mineral dust samples, and therefore
334 put the explanation proposed by Baker and Jickells (2006) into doubt.



335
 336 **Figure 5.** Aerosol Al solubility versus total aerosol Al concentrations: (a) submicron particles
 337 at Xi'an, (b) supermicron particles at Xi'an, (c) submicron particles at Qingdao, (d)
 338 supermicron particles at Qingdao.

339
 340 Aerosol Fe solubility was also frequently observed to increase with the decrease in total
 341 Fe concentrations (Sedwick et al., 2007; Mahowald et al., 2018; Meskhidze et al., 2019), and
 342 one possible reason is the influence of anthropogenic aerosol Fe (Sholkovitz et al., 2009; Ito
 343 and Shi, 2016) with higher solubility than mineral dust (Schroth et al., 2009; Fu et al., 2012;
 344 Ito et al., 2021). Nevertheless, being different from aerosol Fe, aerosol Al stems
 345 predominantly from mineral dust, with little contribution from anthropogenic sources;

346 furthermore, Al solubility was measured to be $0.4\pm 0.6\%$ for coal fly ash (Li et al., 2022), an
347 important type of anthropogenic aerosols, not higher than that for mineral dust ($0.8\pm 0.4\%$).
348 Therefore, we suggest that anthropogenic emission may not be able to explain the inverse
349 dependence of Al aerosol solubility on total Al concentrations.

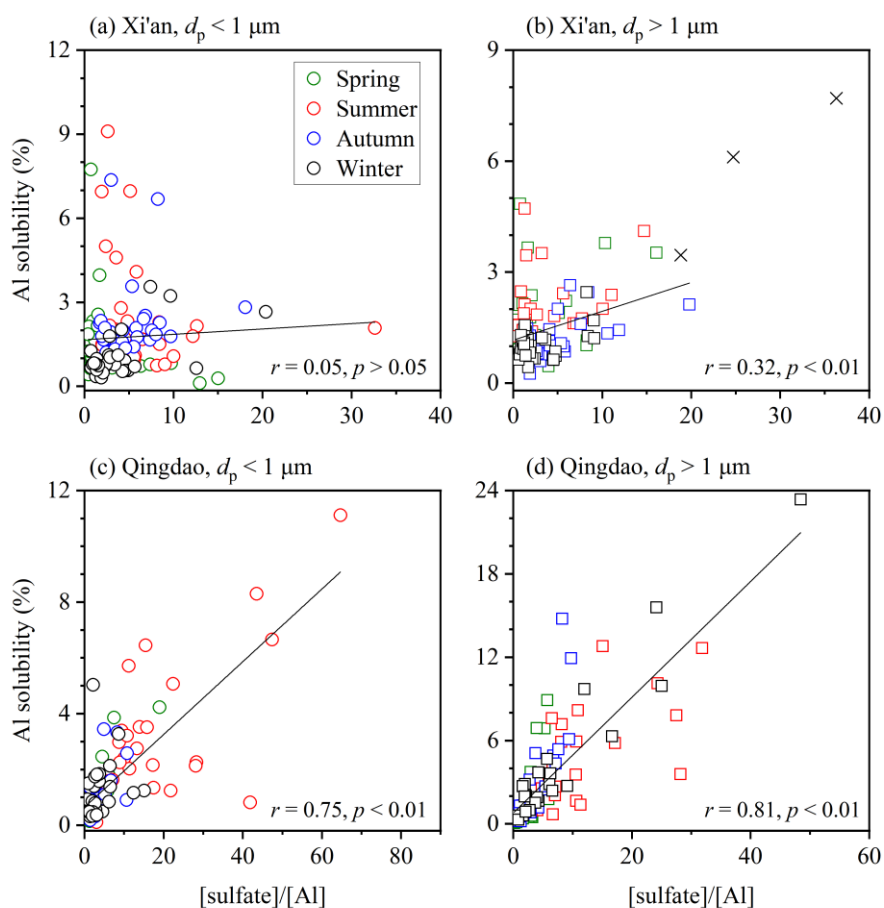
350 We argue that chemical processing in the atmosphere can very well explain such inverse
351 dependence. Total aerosol Al concentrations decrease with transport due to deposition, while
352 reactions with acidic gases (such as SO_2 and NO_x) can enhance the dissolution of aerosol Al
353 (Jickells et al., 2016). Figure 5 shows that the inverse dependence of Al solubility on total Al
354 concentration was more pronounced at Qingdao, with the slopes (b values) much larger than
355 those obtained at Xi'an. This is because compared to Xi'an, Qingdao is more distant from
356 deserts and therefore dust aerosol is expected to be more aged at Qingdao. It also further
357 supports the vital role chemical aging plays in regulating aerosol Al solubility,

358 **4.1 Effects of acid processing and the role of RH**

359 **4.1.1 Effects of acid processing**

360 Laboratory experiments found that the amount of Al dissolved from minerals would
361 increase with the decrease in solution pH (Amram and Ganor, 2005; Bibi et al., 2011, 2014;
362 Cappelli et al., 2018), and some field measurements also suggested that acid processing in the
363 atmosphere could lead to large increase in aerosol Al solubility (Measures et al., 2010; Sakata
364 et al., 2023). In this work, we examined the relationship between aerosol Al solubility and the
365 relative abundance of acidic species ($[\text{sulfate}]/[\text{Al}]$ and $[\text{nitrate}]/[\text{Al}]$) at Xi'an and Qingdao. It
366 should be noted that non-sea-salts sulfate (Virkkula et al., 2006), instead of sulfate, was used
367 at Qingdao because it is a coastal city and heavily impacted by sea spray aerosol.

368 At Xi'an, overall aerosol Al solubility showed no significant correlation with [sulfate]/[Al]
 369 or [nitrate]/[Al] for either supermicron or submicron particles ($r < 0.4$, Figure 6 and S1),
 370 indicating that acid processing did not enhance aerosol Al solubility. Enhancement of aerosol
 371 trace element solubility by acid processing requires internal mixing of acid species with mineral
 372 dust particles (Baker and Croot, 2010). Previous studies suggested that mineral dust particles
 373 observed at Xi'an which is close to deserts largely remained externally mixed with acid species
 374 (Wang et al., 2014; Wu et al., 2017), and thus aerosol Al solubility was not apparently enhanced
 375 by acid processing at Xi'an.
 376



377
 378 **Figure 6.** Aerosol Al solubility versus [sulfate]/[Al]: (a) submicron particles at Xi'an, (b)
 379 supermicron particles at Xi'an, (c) submicron particles at Qingdao, (d) supermicron particles

380 at Qingdao (the r value changed from 0.81 to 0.74 if the data point with the highest Al solubility
381 was excluded). Data represented by crosses are not included in fitting.

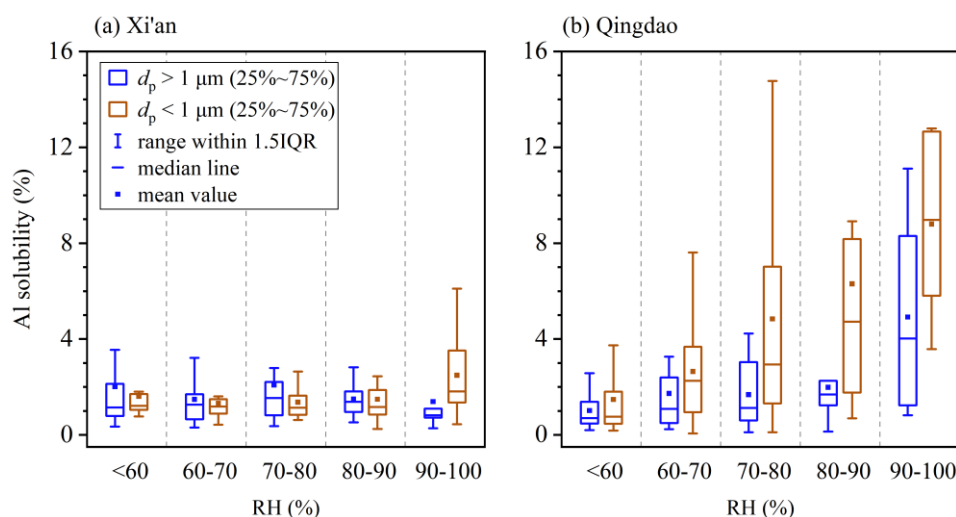
382

383 On the contrary, Figure 6 shows that aerosol Al solubility at Qingdao was well correlated
384 with [sulfate]/[Al] ($r > 0.7$, $p < 0.01$), implying that acidic species were internally mixed with
385 mineral dust particles and thus acid-promoted dissolution significantly enhanced Al solubility.
386 We also found that correlations of Al solubility with [sulfate]/[Al] was better than those with
387 [nitrate]/[Al] (Figures 6 and S1, Table S8), in line with a previous study (Sakata et al., 2023)
388 which found aerosol Al solubility at Hiroshima, southern Japan, to be correlated with
389 [sulfate]/[Al] but not with [nitrate]/[Al]. This may imply that chemical processing by sulfate
390 was more important than nitrate for Al solubility enhancement via acid processing, likely
391 because aluminosilicate dust particles tend to react preferentially with SO_2 and H_2SO_4 while
392 nitrogen oxides react mainly with carbonate particles (Sullivan et al., 2007; Fitzgerald et al.,
393 2015). Furthermore, our work reveals better correlations between Al solubility and [sulfate]/[Al]
394 for supermicron particles than submicron particles (Figure 6), indicating that the effect of acid
395 processing on Al solubility was more important in supermicron particles.

396 4.1.2 The role of RH

397 Relative humidity (RH) is a vital factor influencing liquid water contents and phase state
398 of aerosol particles and thus their secondary chemistry. When RH increased $>60\%$, the phase
399 state of aerosol particles in Northern China changed from semisolid to liquid (Liu et al., 2017;
400 Sun et al., 2018; Song et al., 2022), leading to large increase in aerosol liquid water content and
401 thereby potentially affecting aerosol Al solubility.

402 We observed no apparent variation of aerosol Al solubility with RH at Xi'an (Figure 7a).
 403 When RH was <60%, median Al solubilities for supermicron and submicron particles were
 404 1.22% and 1.14%, respectively; when RH increased >90%, the median Al solubilities were
 405 determined to be 1.82% and 0.82%, showing no obvious increase when compared to those at
 406 <60% RH. This again may imply that chemical processing had very limited impact on aerosol
 407 Al solubility at Xi'an, as mineral dust particles mostly remained externally mixed with
 408 secondary species and their aging extent was very limited (Wang et al., 2014; Wu et al., 2017).



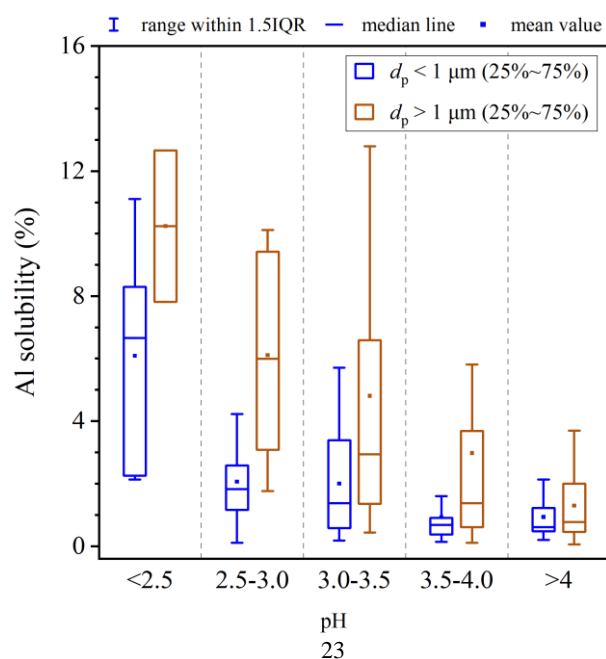
409
 410 **Figure 7.** Aerosol Al solubility at different relative humidity (RH) for submicron and
 411 supermicron particles: (a) Xi'an, (b) Qingdao.

412
 413 In contrast, RH played an important role in regulating aerosol Al solubility at Qingdao,
 414 because mineral dust particles observed at Qingdao had been transported through the North
 415 China Plain and were substantially aged. As shown in Figure 7b, for supermicron particles, the
 416 median Al solubility was only 0.76% at <60% RH, and gradually increased to 4.73% at 80-90%
 417 RH, and abruptly increased to 8.87% at >90% RH. For submicron particles, median Al

418 solubility was <1% at <60% RH, and further increase in RH to 80-90% did not lead to large
419 changes in Al solubility; nevertheless, when RH exceeded 90%, the median Al solubility was
420 remarkably increased to 4.02%, much higher than those observed when RH was < 90%.

421 4.2 Effects of aerosol acidity on aerosol Al solubility at Qingdao

422 Figure 8 shows the dependence of aerosol Al solubility on aerosol acidity (represented by
423 pH) at Qingdao (we did not measure NH₃ at Xi'an and thus could not estimate the aerosol
424 acidity in a reliable manner). For supermicron particles, the median Al solubility was only 0.99%
425 when aerosol pH was >4.0, and gradually increased to 10.24% as aerosol pH was decreased to
426 <2.5. For submicron particles, the median Al solubility was only 0.69% when pH was >4.0,
427 increased slightly with the decrease in pH when pH was in the range of 2.5-4.0, and then
428 increased greatly to 6.09% when pH was decreased to <2.5. In addition, aerosol acidity at
429 Qingdao was highest in summer and lowest in spring (Chen et al., 2024), consistent with the
430 seasonal variation of aerosol Al solubility, further supporting the importance of aerosol acidity
431 in regulating Al solubility.



432

433 **Figure 8.** Aerosol Al solubility corresponding to different aerosol acidity for submicron and
434 supermicron particles in Qingdao.

435

436 As shown in Figure S2, aerosol Al solubility was generally <2% when aerosol acidity was
437 low (pH > 4.0), and higher Al solubility (>2%) was usually observed for samples with high RH
438 and high acidity (pH < 4.0), again underscoring the roles of aerosol acidity (and RH). However,
439 some samples exhibited low Al solubility although the corresponding RH and aerosol acidity
440 were both higher, and such phenomenon was more pronounced for submicron particles. This
441 is very likely linked with aerosol mixing state (Riemer et al., 2019). Aerosol Al solubility and
442 acidity used in our work are both the average properties of an aerosol sample which contains
443 numerous particles, while in reality the two properties will have large particle-to-particle
444 variations. For a given aerosol sample, it can happen that particles with high acidity may
445 contain very little Al while particles with low acidity are enriched in Al; in this case, high
446 acidity do not promote Al solubility for this sample. Single particle analysis which provides
447 mixing state information can give further insights. We also note that samples with low Al
448 solubility but high RH and high acidity were mostly found in clean days, perhaps due to the
449 influence of local resuspended dust for which chemical aging was very limited.

450 **4.3 Size-dependence of aerosol Al solubility**

451 At Xi'an, no obvious difference in aerosol Al solubility was found between supermicron
452 and submicron particles across all the four seasons (Figure 3a). This is because the aging extent
453 of dust particles was rather limited at Xi'an (Wang et al., 2014; Wu et al., 2017) and Al
454 solubility does not vary with particle size for unaged dust particles (Shi et al., 2011). At

455 Qingdao, aerosol Al solubility showed no obvious difference between supermicron and
456 submicron particles in spring, because the aging extent of dust arriving at Qingdao was also
457 limited in spring when Asian dust occurred most frequently. However, in the other three
458 seasons, Al solubility was higher for supermicron particles than submicron particles at Qingdao,
459 and the ratios of median Al solubility in supermicron particles to that in submicron particles
460 were found to be 1.53, 1.70 and 2.57 in summer, autumn and winter, respectively. Similar to
461 our observation at Qingdao, Li et al. (2017) found that aerosol Al solubility was much higher
462 for TSP (14-28%) than PM_{2.5} (2-23%) at the summit of Mount Heng, southern China.

463 On the other hand, a few other studies (Baker et al., 2020; Hsieh et al., 2023; Sakata et al.,
464 2023; Yang et al., 2023) found that aerosol Al solubility was higher in fine particles than coarse
465 particles. For example, aerosol Al solubility was found to increase with the decrease in particle
466 size over the tropical eastern Atlantic (Baker et al., 2020), being ~10.31% for particles in the
467 size of 0.36-0.61 μm and 0.43-4.53% for particles above 0.61 μm . At Hiroshima, southern
468 Japan, aerosol Al solubility was reported to be $8.82\pm 6.48\%$ for fine particles ($<1.3 \mu\text{m}$), more
469 than two times larger than that ($3.25\pm 3.41\%$) for coarse particles ($>1.3 \mu\text{m}$) (Sakata et al., 2023).
470 Baker and Jickells (2006) suggested that this is because fine particles have larger surface-to-
471 volume ratios and thus facilitate Al dissolution via acid processing. Hsieh et al. (2023) found
472 aerosol Al solubility to be 38% for fine particles (0.57-1.0 μm) but only 0.37% for coarse
473 particles ($>7.3 \mu\text{m}$) over the East China Sea, and suggested that the observed size-dependence
474 could be explained by the enrichment of anthropogenic Al (which has higher solubility than
475 dust Al) in fine particles. However, aerosol Al originates predominantly from mineral dust,
476 with little contribution from anthropogenic sources (Taylor and McLennan, 1985; Mahowald

477 et al., 2018), and fractional solubility of anthropogenic Al was not necessarily higher than
478 mineral dust (Li et al., 2022).

479 As discussed above, there is not clear yet how and why aerosol Al solubility varies with
480 particle size. Such discrepancy is at least partly because different leaching protocols were used
481 in previous studies to extract dissolved aerosol Al and thereby Al solubility obtained in
482 different studies was not directly comparable (Meskhidze et al., 2019; Li et al., 2023; Li et al.,
483 2024). Furthermore, mechanistic insights can be obtained by laboratory experiments which
484 examine the size dependence of the solubility and dissolution kinetics of Al for mineral dust
485 particles under atmospherically relevant conditions.

486 **5. Conclusions and atmospheric implications**

487 Deposition of mineral dust aerosol is a major external source of several nutrient and toxic
488 elements for surface water in open oceans, and thus have large impacts on marine
489 biogeochemistry; however, previous studies which estimated dust deposition flux into the
490 oceans reveals large discrepancies. Aerosol Al solubility, which is a critical parameter in using
491 dissolved Al concentrations in surface seawater as a tracer to constrain dust deposition flux,
492 remains poorly understood. In this work, we investigated seasonal variations of aerosol Al
493 solubility for supermicron ($>1 \mu\text{m}$) and submicron ($<1 \mu\text{m}$) aerosol particles at Xi'an and
494 Qingdao, both located in Northern China, in attempt to elucidate the processes and mechanisms
495 which govern the variation of aerosol Al solubility in the atmosphere.

496 At Xi'an, aerosol Al solubility was low in general for both supermicron and submicron
497 particles, showing no obvious variability in different seasons or under different weather
498 conditions. This implies that chemical processing did not substantially enhance aerosol Al

499 solubility at Xi'an, as it is an inland city close to major deserts in Northwestern China and thus
500 the aging extent of mineral dust particles arriving at Xi'an was quite limited. Compared to
501 Xi'an, aerosol Al solubility was higher at Qingdao, a coastal city in Northern China;
502 furthermore, Al solubility was higher in the other three seasons than in spring, and much higher
503 for haze- and especially fog-impacted days than dust days. This indicates that chemical
504 processing substantially increased aerosol Al solubility at Qingdao.

505 Aerosol Al solubility at Xi'an showed no significant correlation with relative abundance
506 of sulfate or nitrate, and did not vary apparently with RH; in contrast, Al solubility at Qingdao
507 was well correlated with relative abundance of sulfate and nitrate, and increased with RH. This
508 further supports that chemical processing had little impact on aerosol Al solubility at Xi'an
509 (because the aging extent of mineral dust aerosol at Xi'an is very limited) but remarkably
510 increased aerosol Al solubility at Qingdao (because mineral dust particles transported to
511 Qingdao were substantially aged). Moreover, for both supermicron and submicron particles,
512 Al solubility at Qingdao was found to increase with aerosol acidity (in addition to RH),
513 underscoring the vital role of aerosol liquid water and acidity in enhancing Al dissolution via
514 chemical aging.

515 Our comprehensive investigation of aerosol Al solubility at two locations in Northern
516 China suggests that atmospheric chemical processing dictates aerosol Al solubility. As a result,
517 aerosol Al solubility is expected to spatially variable, depending on the extent of chemical
518 processing. For example, we found that aerosol Al solubility is higher at Qingdao than Xi'an
519 in general, and expect it to increase further as mineral dust aerosol is further transported
520 eastward to the Pacific. **Although our measurements were only conducted at two sites, our work**

521 provides important insights into processes driving spatiotemporal variability of aerosol AI
522 solubility, and such understanding can aid us to develop aerosol AI solubility parameterizations.

523

524 **Author contribution.**

525 **TZ:** Formal analysis, Investigation, Writing - Original Draft, Writing - Review & Editing;

526 **YC:** Formal analysis, Investigation, Writing - Original Draft; **HZ:** Investigation; **Lei Liu:**

527 Writing - Review & Editing; **CH:** Investigation; **ZF:** Investigation; **YZ:** Investigation; **FW:**

528 Resources; **Lan Luo:** Resources; **GZ:** Writing - Review & Editing; **XW:** Resources; **MT:**

529 Conceptualization, Formal analysis, Supervision; Writing - Original Draft, Writing - Review

530 & Editing.

531 **Competing interests.**

532 The authors declare that they have no conflict of interest.

533 **Acknowledgement.**

534 We would like to thank colleagues at Shandong University, Shaanxi University of Science

535 and Technology, and Institute of Earth Environment, Chinese Academy of Sciences for their

536 support during field measurements.

537 **Financial support.**

538 This work was sponsored by National Natural Science Foundation of China (42277088,

539 42407149 and 22361162668), Guangzhou Bureau of Science and Technology

540 (2024A04J6533), International Partnership Program of Chinese Academy of Sciences

541 (164GJHZ2024011FN), Guangdong Basic and Applied Basic Research Fund Committee

542 (2023A1515012010), and Guangdong Foundation for Program of Science and Technology

543 Research (2023B1212060049).

544

545

546 **References**

- 547 Aghnatiou, C., Losno, R., and Dulac, F.: A fine fraction of soil used as an aerosol analogue during the DUNE
548 experiment: sequential solubility in water, decreasing pH step-by-step, *Biogeosciences*, 11, 4627-4633,
549 <https://doi.org/10.5194/bg-11-4627-2014>, 2014.
- 550 Amram, K., and Ganor, J.: The combined effect of pH and temperature on smectite dissolution rate under acidic
551 conditions, *Geochim. Cosmochim. Acta*, 69, 2535-2546, <https://doi.org/10.1016/j.gca.2004.10.001>,
552 2005.
- 553 An, Z., Huang, R., Zhang, R., Tie, X., Li, G., Cao, J., Zhou, W., Shi, Z., Han, Y., Gu, Z., and Ji, Y.: Severe haze
554 in northern China: A synergy of anthropogenic emissions and atmospheric processes, *PNAS*, 116,
555 8657-8666, <https://doi.org/10.1073/pnas.1900125116>, 2019.
- 556 Anderson, R. F., Cheng, H., Edwards, R. L., Fleisher, M. Q., Hayes, C. T., Huang, K. F., Kadko, D., Lam, P. J.,
557 Landing, W. M., Lao, Y., Lu, Y., Measures, C. I., Moran, S. B., Morton, P. L., Ohnemus, D. C.,
558 Robinson, L. F., and Shelley, R. U.: How well can we quantify dust deposition to the ocean?, *Phil.
559 Trans. R. Soc. A*, 374: 20150285, <https://doi.org/10.1098/rsta.2015.0285>, 2016.
- 560 Baker, A. R., and Jickells, T. D.: Mineral particle size as a control on aerosol iron solubility, *Geophys. Res. Lett.*,
561 33, L17608, <https://doi.org/10.1029/2006gl026557>, 2006.
- 562 Baker, A. R., Jickells, T. D., Witt, M., and Linge, K. L.: Trends in the solubility of iron, aluminium, manganese
563 and phosphorus in aerosol collected over the Atlantic Ocean, *Mar. Chem.*, 98, 43-58,
564 <https://doi.org/10.1016/j.marchem.2005.06.004>, 2006.
- 565 Baker, A. R., and Croot, P. L.: Atmospheric and marine controls on aerosol iron solubility in seawater, *Mar.
566 Chem.*, 120, 4-13, <https://doi.org/10.1016/j.marchem.2008.09.003>, 2010.
- 567 Baker, A. R., Li, M., and Chance, R.: Trace Metal Fractional Solubility in Size - Segregated Aerosols From the
568 Tropical Eastern Atlantic Ocean, *Global Biogeochem. Cycles*, 34, e2019GB006510,
569 <https://doi.org/10.1029/2019gb006510>, 2020.
- 570 Benaltabet, T., Lapid, G., and Torfstein, A.: Dissolved aluminium dynamics in response to dust storms, wet
571 deposition, and sediment resuspension in the Gulf of Aqaba, northern Red Sea, *Geochim. Cosmochim.
572 Acta*, 335, 137-154, <https://doi.org/10.1016/j.gca.2022.08.029>, 2022.
- 573 Bibi, I., Singh, B., and Silvester, E.: Dissolution of illite in saline-acidic solutions at 25°C, *Geochim.
574 Cosmochim. Acta*, 75, 3237-3249, <https://doi.org/10.1016/j.gca.2011.03.022>, 2011.
- 575 Bibi, I., Singh, B., and Silvester, E.: Dissolution kinetics of soil clays in sulfuric acid solutions: Ionic strength
576 and temperature effects, *Appl. Geochem.*, 51, 170-183,
577 <https://doi.org/10.1016/j.apgeochem.2014.10.004>, 2014.
- 578 Buck, C. S., Landing, W. M., Resing, J. A., and Measures, C. I.: The solubility and deposition of aerosol Fe and
579 other trace elements in the North Atlantic Ocean: Observations from the A16N CLIVAR/CO2 repeat
580 hydrography section, *Mar. Chem.*, 120, 57-70, <https://doi.org/10.1016/j.marchem.2008.08.003>, 2010.
- 581 Buck, C. S., Landing, W. M., and Resing, J.: Pacific Ocean aerosols: Deposition and solubility of iron,
582 aluminum, and other trace elements, *Mar. Chem.*, 157, 117-130,
583 <https://doi.org/10.1016/j.marchem.2013.09.005>, 2013.
- 584 Cai, Q.-L., Dai, X.-R., Li, J.-R., Tong, L., Hui, Y., Cao, M.-Y., Li, M., and Xiao, H.: The characteristics and
585 mixing states of PM2.5 during a winter dust storm in Ningbo of the Yangtze River Delta, China, *Sci.
586 Total Environ.*, 709, 136146, <https://doi.org/10.1016/j.scitotenv.2019.136146>, 2020.
- 587 Cao, J. J., Chow, J. C., Watson, J. G., Wu, F., Han, Y. M., Jin, Z. D., Shen, Z. X., and An, Z. S.: Size-

588 differentiated source profiles for fugitive dust in the Chinese Loess Plateau, *Atmos. Environ.*, 42, 2261-
589 2275, <https://doi.org/10.1016/j.atmosenv.2007.12.041>, 2008.

590 Cao, J. J., and Cui, L.: Current Status, Characteristics and Causes of Particulate Air Pollution in the Fenwei
591 Plain, China: A Review, *J. Geophys. Res.-Atmos.*, 126, e2020JD034472,
592 <https://doi.org/10.1029/2020JD034472>, 2021.

593 Cappelli, C., Yokoyama, S., Cama, J., and Huertas, F. J.: Montmorillonite dissolution kinetics: Experimental and
594 reactive transport modeling interpretation, *Geochim. Cosmochim. Acta*, 227, 96-122,
595 <https://doi.org/10.1016/j.gca.2018.01.039>, 2018.

596 Chance, R., Jickells, T. D., and Baker, A. R.: Atmospheric trace metal concentrations, solubility and deposition
597 fluxes in remote marine air over the south-east Atlantic, *Mar. Chem.*, 177, 45-56,
598 <https://doi.org/10.1016/j.marchem.2015.06.028>, 2015.

599 Chen, Y., Wang, Z., Fang, Z., Huang, C., Xu, H., Zhang, H., Zhang, T., Wang, F., Luo, L., Shi, G., Wang, X., and
600 Tang, M.: Dominant Contribution of Non-dust Primary Emissions and Secondary Processes to
601 Dissolved Aerosol Iron, *Environ. Sci. Technol.*, 58, 17355-17363,
602 <https://doi.org/10.1021/acs.est.4c05816>, 2024.

603 Duvall, R. M., Majestic, B. J., Shafer, M. M., Chuang, P. Y., Simoneit, B. R. T., and Schauer, J. J.: The water-
604 soluble fraction of carbon, sulfur, and crustal elements in Asian aerosols and Asian soils, *Atmos.*
605 *Environ.*, 42, 5872-5884, <https://doi.org/10.1016/j.atmosenv.2008.03.028>, 2008.

606 Fang, Z., Dong, S., Huang, C., Jia, S., Wang, F., Liu, H., Meng, H., Luo, L., Chen, Y., Zhang, H., Li, R., Zhu, Y.,
607 and Tang, M.: On using an aerosol thermodynamic model to calculate aerosol acidity of coarse
608 particles, *J. Environ. Sci.*, 148, 46-56, <https://doi.org/10.1016/j.jes.2023.07.001>, 2025.

609 Fitzgerald, E., Ault, A. P., Zauscher, M. D., Mayol-Bracero, O. L., and Prather, K. A.: Comparison of the mixing
610 state of long-range transported Asian and African mineral dust, *Atmos. Environ.*, 115, 19-25,
611 <https://doi.org/10.1016/j.atmosenv.2015.04.031>, 2015.

612 Fountoukis, C., and Nenes, A.: ISORROPIA II: a computationally efficient thermodynamic equilibrium model
613 for K^+ - Ca^{2+} - Mg^{2+} - NH_4^+ - Na^+ - SO_4^{2-} - NO_3^- - Cl^- - H_2O aerosols, *Atmos. Chem. Phys.*, 7, 4639-4659,
614 <https://doi.org/10.5194/acp-7-4639-2007>, 2007.

615 Fu, H., Lin, J., Shang, G., Dong, W., Grassian, V. H., Carmichael, G. R., Li, Y., and Chen, J.: Solubility of Iron
616 from Combustion Source Particles in Acidic Media Linked to Iron Speciation, *Environ. Sci. Technol.*,
617 46, 11119-11127, <https://doi.org/10.1021/es302558m>, 2012.

618 Grand, M. M., Measures, C. I., Hata, M., Hiscock, W. T., Buck, C. S., and Landing, W. M.: Dust deposition in
619 the eastern Indian Ocean: The ocean perspective from Antarctica to the Bay of Bengal, *Global*
620 *Biogeochem. Cycles*, 29, 357-374, <https://doi.org/10.1002/2014GB004898>, 2015.

621 Guo, L., Chen, Y., Wang, F., Meng, X., Xu, Z., and Zhuang, G.: Effects of Asian dust on the atmospheric input
622 of trace elements to the East China Sea, *Mar. Chem.*, 163, 19-27,
623 <https://doi.org/10.1016/j.marchem.2014.04.003>, 2014.

624 Han, Q., Moore, J. K., Zender, C., Measures, C., and Hydes, D.: Constraining oceanic dust deposition using
625 surface ocean dissolved Al, *Global Biogeochem. Cycles*, 22, GB2003,
626 <https://doi.org/10.1029/2007GB002975>, 2008.

627 Haugvaldstad, O. W., Tang, H., Kaakinen, A., Bohm, K., Groot Zwaafink, C. D., Grythe, H., Stevens, T.,
628 Zhang, Z., and Stordal, F.: Spatial Source Contribution and Interannual Variation in Deposition of Dust
629 Aerosols Over the Chinese Loess Plateau, *J. Geophys. Res.-Atmos.*, 129, e2023JD040470,
630 <https://doi.org/10.1029/2023JD040470>, 2024.

631 Hsieh, C.-C., You, C.-F., and Ho, T.-Y.: The solubility and deposition flux of East Asian aerosol metals in the
632 East China Sea: The effects of aeolian transport processes, *Mar. Chem.*, 253, 104268,
633 <https://doi.org/10.1016/j.marchem.2023.104268>, 2023.

634 Hsu, S.-C., Wong, G. T. F., Gong, G.-C., Shiah, F.-K., Huang, Y.-T., Kao, S.-J., Tsai, F., Candice Lung, S.-C.,
635 Lin, F.-J., Lin, I. I., Hung, C.-C., and Tseng, C.-M.: Sources, solubility, and dry deposition of aerosol
636 trace elements over the East China Sea, *Mar. Chem.*, 120, 116-127,
637 <https://doi.org/10.1016/j.marchem.2008.10.003>, 2010.

638 Huang, R.-J., Zhang, Y., Bozzetti, C., Ho, K.-F., Cao, J.-J., Han, Y., Daellenbach, K. R., Slowik, J. G., Platt, S.
639 M., Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Brun, E. A., Crippa, M., Ciarelli, G., Piazzalunga,
640 A., Schwikowski, M., Abbaszade, G., Schnelle-Kreis, J., Zimmermann, R., An, Z., Szidat, S.,
641 Baltensperger, U., Haddad, I. E., and Prévôt, A. S. H.: High secondary aerosol contribution to
642 particulate pollution during haze events in China, *Nature*, 514, 218-222, [10.1038/nature13774](https://doi.org/10.1038/nature13774), 2014.

643 Huneus, N., Schulz, M., Balkanski, Y., Griesfeller, J., Prospero, J., Kinne, S., Bauer, S., Boucher, O., Chin, M.,
644 Dentener, F., Diehl, T., Easter, R., Fillmore, D., Ghan, S., Ginoux, P., Grini, A., Horowitz, L., Koch, D.,
645 Krol, M. C., Landing, W., Liu, X., Mahowald, N., Miller, R., Morcrette, J. J., Myhre, G., Penner, J.,
646 Perlwitz, J., Stier, P., Takemura, T., and Zender, C. S.: Global dust model intercomparison in AeroCom
647 phase I, *Atmos. Chem. Phys.*, 11, 7781-7816, <https://doi.org/10.5194/acp-11-7781-2011>, 2011.

648 Ito, A., and Shi, Z.: Delivery of anthropogenic bioavailable iron from mineral dust and combustion aerosols to
649 the ocean, *Atmos. Chem. Phys.*, 16, 85-99, <https://doi.org/10.5194/acp-16-85-2016>, 2016.

650 Ito, A., Ye, Y., Baldo, C., and Shi, Z.: Ocean fertilization by pyrogenic aerosol iron, *npj Clim. Atmos. Sci.*, 4, 30,
651 <https://doi.org/10.1038/s41612-021-00185-8>, 2021.

652 Jeong, G. Y.: Mineralogy and geochemistry of Asian dust: dependence on migration path, fractionation, and
653 reactions with polluted air, *Atmos. Chem. Phys.*, 20, 7411-7428, [https://doi.org/10.5194/acp-20-7411-](https://doi.org/10.5194/acp-20-7411-2020)
654 2020, 2020.

655 Jiang, H.-B., Hutchins, D. A., Zhang, H.-R., Feng, Y.-Y., Zhang, R.-F., Sun, W.-W., Ma, W., Bai, Y., Wells, M.,
656 He, D., Jiao, N., Wang, Y., and Chai, F.: Complexities of regulating climate by promoting marine
657 primary production with ocean iron fertilization, *Earth Sci. Rev.*, 249, 104675,
658 <https://doi.org/10.1016/j.earscirev.2024.104675>, 2024.

659 Jickells, T. D., An, Z. S., Andersen, K. K., Baker, A. R., Bergametti, G., Brooks, N., Cao, J. J., Boyd, P. W.,
660 Duce, R. A., Hunter, K. A., Kawahata, H., Kubilay, N., laRoche, J., Liss, P. S., Mahowald, N.,
661 Prospero, J. M., Ridgwell, A. J., Tegen, I., and Torres, R.: Global Iron Connections Between Desert
662 Dust, Ocean Biogeochemistry, and Climate, *Science*, 308, 67-71,
663 <https://doi.org/10.1126/science.1105959>, 2005.

664 Jickells, T. D., Baker, A. R., and Chance, R.: Atmospheric transport of trace elements and nutrients to the
665 oceans, *Phil. Trans. R. Soc. A*, 374, 20150286, <https://doi.org/10.1098/rsta.2015.0286>, 2016.

666 Kok, J. F., Storelvmo, T., Karydis, V. A., Adebisi, A. A., Mahowald, N. M., Evan, A. T., He, C., and Leung, D.
667 M.: Mineral dust aerosol impacts on global climate and climate change, *Nat. Rev. Earth Environ.*, 4,
668 71-86, <https://doi.org/10.1038/s43017-022-00379-5>, 2023.

669 Li, R., Zhang, H., Wang, F., Ren, Y., Jia, S., Jiang, B., Jia, X., Tang, Y., and Tang, M.: Abundance and fractional
670 solubility of phosphorus and trace metals in combustion ash and desert dust: Implications for
671 bioavailability and reactivity, *Sci. Total Environ.*, 816, 151495,
672 <https://doi.org/10.1016/j.scitotenv.2021.151495>, 2022.

673 Li, R., Dong, S., Huang, C., Yu, F., Wang, F., Li, X., Zhang, H., Ren, Y., Guo, M., Chen, Q., Ge, B., and Tang,

674 M.: Evaluating the effects of contact time and leaching solution on measured solubilities of aerosol
675 trace metals, *Appl. Geochem.*, 148, 105551, <https://doi.org/10.1016/j.apgeochem.2022.105551>, 2023.

676 Li, R., Panda, P. P., Chen, Y., Zhu, Z., Wang, F., Zhu, Y., Meng, H., Ren, Y., Kumar, A., and Tang, M.: Aerosol
677 trace element solubility determined using ultrapure water batch leaching: an intercomparison study of
678 four different leaching protocols, *Atmos. Meas. Tech.*, 17, 3147-3156, [https://doi.org/10.5194/amt-17-](https://doi.org/10.5194/amt-17-3147-2024)
679 3147-2024, 2024.

680 Li, T., Wang, Y., Zhou, J., Wang, T., Ding, A., Nie, W., Xue, L., Wang, X., and Wang, W.: Evolution of trace
681 elements in the planetary boundary layer in southern China: Effects of dust storms and aerosol-cloud
682 interactions, *J. Geophys. Res.-Atmos.*, 122, 3492-3506, <https://doi.org/10.1002/2016JD025541>, 2017.

683 Li, W., Shao, L., Shi, Z., Chen, J., Yang, L., Yuan, Q., Yan, C., Zhang, X., Wang, Y., Sun, J., Zhang, Y., Shen, X.,
684 Wang, Z., and Wang, W.: Mixing state and hygroscopicity of dust and haze particles before leaving
685 Asian continent, *J. Geophys. Res.-Atmos.*, 119, 1044-1059, <https://doi.org/10.1002/2013JD021003>,
686 2014.

687 Liu, Y., Wu, Z., Wang, Y., Xiao, Y., Gu, F., Zheng, J., Tan, T., Shang, D., Wu, Y., Zeng, L., Hu, M., Bateman, A.
688 P., and Martin, S. T.: Submicrometer Particles Are in the Liquid State during Heavy Haze Episodes in
689 the Urban Atmosphere of Beijing, China, *Environ. Sci. Technol. Lett.*, 4, 427-432,
690 <https://doi.org/10.1021/acs.estlett.7b00352>, 2017.

691 López-García, P., Gelado-Caballero, M. D., Collado-Sánchez, C., and Hernández-Brito, J. J.: Solubility of
692 aerosol trace elements: Sources and deposition fluxes in the Canary Region, *Atmos. Environ.*, 148,
693 167-174, <https://doi.org/10.1016/j.atmosenv.2016.10.035>, 2017.

694 Mahowald, N.: Aerosol Indirect Effect on Biogeochemical Cycles and Climate, *Science*, 334, 794-796,
695 <https://doi.org/10.1126/science.1207374>, 2011.

696 Mahowald, N. M., Hamilton, D. S., Mackey, K. R. M., Moore, J. K., Baker, A. R., Scanza, R. A., and Zhang, Y.:
697 Aerosol trace metal leaching and impacts on marine microorganisms, *Nat. Commun.*, 9, 2614,
698 <https://doi.org/10.1038/s41467-018-04970-7>, 2018.

699 Measures, C. I., and Brown, E. T.: Estimating Dust Input to the Atlantic Ocean Using Surface Water Aluminium
700 Concentrations, in: *The Impact of Desert Dust Across the Mediterranean*, edited by: Guerzoni, S., and
701 Chester, R., Springer Netherlands, Dordrecht, 301-311, 1996.

702 Measures, C. I., and Vink, S.: On the use of dissolved aluminum in surface waters to estimate dust deposition to
703 the ocean, *Global Biogeochem. Cycles*, 14, 317-327, <https://doi.org/10.1029/1999GB001188>, 2000.

704 Measures, C. I., Sato, T., Vink, S., Howell, S., and Li, Y. H.: The fractional solubility of aluminium from mineral
705 aerosols collected in Hawaii and implications for atmospheric deposition of biogeochemically
706 important trace elements, *Mar. Chem.*, 120, 144-153, <https://doi.org/10.1016/j.marchem.2009.01.014>,
707 2010.

708 Meskhidze, N., Völker, C., Al-Abadleh, H. A., Barbeau, K., Bressac, M., Buck, C., Bundy, R. M., Croot, P.,
709 Feng, Y., Ito, A., Johansen, A. M., Landing, W. M., Mao, J., Myriokefalitakis, S., Ohnemus, D.,
710 Pasquier, B., and Ye, Y.: Perspective on identifying and characterizing the processes controlling iron
711 speciation and residence time at the atmosphere-ocean interface, *Mar. Chem.*, 217, 103704,
712 <https://doi.org/10.1016/j.marchem.2019.103704>, 2019.

713 Moore, C. M., Mills, M. M., Arrigo, K. R., Berman-Frank, I., Bopp, L., Boyd, P. W., Galbraith, E. D., Geider, R.
714 J., Guieu, C., Jaccard, S. L., Jickells, T. D., La Roche, J., Lenton, T. M., Mahowald, N. M., Marañón,
715 E., Marinov, I., Moore, J. K., Nakatsuka, T., Oeschlies, A., Saito, M. A., Thingstad, T. F., Tsuda, A., and
716 Ulloa, O.: Processes and patterns of oceanic nutrient limitation, *Nature Geoscience*, 6, 701-710,

717 <https://doi.org/10.1038/ngeo1765>, 2013.

718 Mulder, J., van Breemen, N., and Eijck, H. C.: Depletion of soil aluminium by acid deposition and implications
719 for acid neutralization, *Nature*, 337, 247-249, 10.1038/337247a0, 1989.

720 Pan, X., Uno, I., Wang, Z., Nishizawa, T., Sugimoto, N., Yamamoto, S., Kobayashi, H., Sun, Y., Fu, P., Tang, X.,
721 and Wang, Z.: Real-time observational evidence of changing Asian dust morphology with the mixing of
722 heavy anthropogenic pollution, *Sci. Rep.*, 7, 335, <https://doi.org/10.1038/s41598-017-00444-w>, 2017.

723 Paris, R., Desboeufs, K. V., Formenti, P., Nava, S., and Chou, C.: Chemical characterisation of iron in dust and
724 biomass burning aerosols during AMMA-SOP0/DABEX: implication for iron solubility, *Atmos. Chem.*
725 *Phys.*, 10, 4273-4282, <https://doi.org/10.5194/acp-10-4273-2010>, 2010.

726 Prospero, J. M., Ginoux, P., Torres, O., Nicholson, S. E., and Gill, T. E.: Environmental characterization of
727 global sources of atmospheric soil dust identified with the Nimbus 7 Total Ozone Mapping
728 Spectrometer (TOMS) absorbing aerosol product, *Rev. Geophys.*, 40, 2-1-2-31,
729 <https://doi.org/10.1029/2000RG000095>, 2002.

730 Riemer, N., Ault, A. P., West, M., Craig, R. L., and Curtis, J. H.: Aerosol Mixing State: Measurements,
731 Modeling, and Impacts, *Rev. Geophys.*, 57, 187-249, <https://doi.org/10.1029/2018RG000615>, 2019.

732 Sakata, K., Sakaguchi, A., Yamakawa, Y., Miyamoto, C., Kurisu, M., and Takahashi, Y.: Measurement report:
733 Stoichiometry of dissolved iron and aluminum as an indicator of the factors controlling the fractional
734 solubility of aerosol iron – results of the annual observations of size-fractionated aerosol particles in
735 Japan, *Atmos. Chem. Phys.*, 23, 9815-9836, <https://doi.org/10.5194/acp-23-9815-2023>, 2023.

736 Schroth, A. W., Crusius, J., Sholkovitz, E. R., and Bostick, B. C.: Iron solubility driven by speciation in dust
737 sources to the ocean, *Nature Geoscience*, 2, 337-340, <https://doi.org/10.1038/ngeo501>, 2009.

738 Schulz, M., Prospero, J. M., Baker, A. R., Dentener, F., Ickes, L., Liss, P. S., Mahowald, N. M., Nickovic, S.,
739 García-Pando, C. P., Rodríguez, S., Sarin, M., Tegen, I., and Duce, R. A.: Atmospheric Transport and
740 Deposition of Mineral Dust to the Ocean: Implications for Research Needs, *Environ. Sci. Technol.*, 46,
741 10390-10404, <https://doi.org/10.1021/es300073u>, 2012.

742 Sedwick, P. N., Sholkovitz, E. R., and Church, T. M.: Impact of anthropogenic combustion emissions on the
743 fractional solubility of aerosol iron: Evidence from the Sargasso Sea, *Geochem. Geophys. Geosyst.*, 8,
744 <https://doi.org/10.1029/2007GC001586>, 2007.

745 Shang, T., Kong, L., and Qi, J.: Metal elements in atmospheric aerosols during different pollution events in the
746 coastal region of the Yellow Sea: Concentration, solubility and deposition flux, *Mar. Pollut. Bull.*, 206,
747 116711, <https://doi.org/10.1016/j.marpolbul.2024.116711>, 2024.

748 Shelley, R. U., Landing, W. M., Ussher, S. J., Planquette, H., and Sarthou, G.: Regional trends in the fractional
749 solubility of Fe and other metals from North Atlantic aerosols (GEOTRACES cruises GA01 and
750 GA03) following a two-stage leach, *Biogeosciences*, 15, 2271-2288, <https://doi.org/10.5194/bg-15-2271-2018>, 2018.

751 Shelley, R. U., Baker, A. R., Thomas, M., and Murphy, S.: Aerosol trace element solubility and deposition fluxes
752 over the Mediterranean Sea and Black Sea basins, *Biogeosciences*, 22, 585-600,
753 <https://doi.org/10.5194/bg-22-585-2025>, 2025.

754 Shi, J., Guan, Y., Ito, A., Gao, H., Yao, X., Baker, A. R., and Zhang, D.: High Production of Soluble Iron
755 Promoted by Aerosol Acidification in Fog, *Geophys. Res. Lett.*, 47, e2019GL086124,
756 <https://doi.org/10.1029/2019GL086124>, 2020.

757 Shi, Z. B., Woodhouse, M. T., Carslaw, K. S., Krom, M. D., Mann, G. W., Baker, A. R., Savov, I., Fones, G. R.,
758 Brooks, B., Drake, N., Jickells, T. D., and Benning, L. G.: Minor effect of physical size sorting on iron
759

760 solubility of transported mineral dust, *Atmos. Chem. Phys.*, 11, 8459-8469, <https://doi.org/10.5194/acp->
761 11-8459-2011, 2011.

762 Sholkovitz, E. R., Sedwick, P. N., and Church, T. M.: Influence of anthropogenic combustion emissions on the
763 deposition of soluble aerosol iron to the ocean: Empirical estimates for island sites in the North
764 Atlantic, *Geochim. Cosmochim. Acta*, 73, 3981-4003, <https://doi.org/10.1016/j.gca.2009.04.029>, 2009.

765 Song, M., Jeong, R., Kim, D., Qiu, Y., Meng, X., Wu, Z., Zuend, A., Ha, Y., Kim, C., Kim, H., Gaikwad, S.,
766 Jang, K. S., Lee, J. Y., and Ahn, J.: Comparison of Phase States of PM(2.5) over Megacities, Seoul and
767 Beijing, and Their Implications on Particle Size Distribution, *Environ. Sci. Technol.*, 56, 17581-17590,
768 <https://doi.org/10.1021/acs.est.2c06377>, 2022.

769 Sullivan, R. C., Guazzotti, S. A., Sodeman, D. A., and Prather, K. A.: Direct observations of the atmospheric
770 processing of Asian mineral dust, *Atmos. Chem. Phys.*, 7, 1213-1236, <https://doi.org/10.5194/acp-7->
771 1213-2007, 2007.

772 Sun, J., Liu, L., Xu, L., Wang, Y., Wu, Z., Hu, M., Shi, Z., Li, Y., Zhang, X., Chen, J., and Li, W.: Key Role of
773 Nitrate in Phase Transitions of Urban Particles: Implications of Important Reactive Surfaces for
774 Secondary Aerosol Formation, *J. Geophys. Res.-Atmos.*, 123, 1234-1243,
775 <https://doi.org/10.1002/2017JD027264>, 2018.

776 Takahashi, Y., Higashi, M., Furukawa, T., and Mitsunobu, S.: Change of iron species and iron solubility in Asian
777 dust during the long-range transport from western China to Japan, *Atmos. Chem. Phys.*, 11, 11237-
778 11252, <https://doi.org/10.5194/acp-11-11237-2011>, 2011.

779 Tang, M., Cziczo, D. J., and Grassian, V. H.: Interactions of Water with Mineral Dust Aerosol: Water
780 Adsorption, Hygroscopicity, Cloud Condensation, and Ice Nucleation, *Chem. Rev.*, 116, 4205-4259,
781 <https://doi.org/10.1021/acs.chemrev.5b00529>, 2016.

782 Taylor, S. R., and McLennan, S. M.: *The continental crust: Its composition and evolution*, Blackwell Scientific
783 Publications, Oxford, 312 pp., 1985.

784 Trochkin, D., Iwasaka, Y., Matsuki, A., Yamada, M., Kim, Y.-S., Nagatani, T., Zhang, D., Shi, G.-Y., and Shen,
785 Z.: Mineral aerosol particles collected in Dunhuang, China, and their comparison with chemically
786 modified particles collected over Japan, *J. Geophys. Res.-Atmos.*, 108, 8642,
787 <https://doi.org/10.1029/2002JD003268>, 2003.

788 Virkkula, A., Teinilä, K., Hillamo, R., Kerminen, V. M., Saarikoski, S., Aurela, M., Viidanoja, J., Paatero, J.,
789 Koponen, I. K., and Kulmala, M.: Chemical composition of boundary layer aerosol over the Atlantic
790 Ocean and at an Antarctic site, *Atmos. Chem. Phys.*, 6, 3407-3421, <https://doi.org/10.5194/acp-6-3407->
791 2006, 2006.

792 Walters, W. W., and Hastings, M. G.: Collection of Ammonia for High Time-Resolved Nitrogen Isotopic
793 Characterization Utilizing an Acid-Coated Honeycomb Denuder, *Anal. Chem.*, 90, 8051-8057,
794 <https://doi.org/10.1021/acs.analchem.8b01007>, 2018.

795 Wang, G. H., Cheng, C. L., Huang, Y., Tao, J., Ren, Y. Q., Wu, F., Meng, J. J., Li, J. J., Cheng, Y. T., Cao, J. J.,
796 Liu, S. X., Zhang, T., Zhang, R., and Chen, Y. B.: Evolution of aerosol chemistry in Xi'an, inland
797 China, during the dust storm period of 2013 - Part 1: Sources, chemical forms and formation
798 mechanisms of nitrate and sulfate, *Atmos. Chem. Phys.*, 14, 11571-11585, <https://doi.org/10.5194/acp->
799 14-11571-2014, 2014.

800 Wang, S., Yan, Q., Zhang, R., Jiang, N., Yin, S., and Ye, H.: Size-fractionated particulate elements in an inland
801 city of China: Deposition flux in human respiratory, health risks, source apportionment, and dry
802 deposition, *Environmental Pollution*, 247, 515-523, <https://doi.org/10.1016/j.envpol.2019.01.051>,

803 2019.

804 Wang, Z., Liu, C., Xie, Z., Hu, Q., Andreae, M. O., Dong, Y., Zhao, C., Liu, T., Zhu, Y., Liu, H., Xing, C., Tan,
805 W., Ji, X., Lin, J., and Liu, J.: Elevated dust layers inhibit dissipation of heavy anthropogenic surface
806 air pollution, *Atmos. Chem. Phys.*, 20, 14917-14932, 10.5194/acp-20-14917-2020, 2020.

807 Westberry, T. K., Behrenfeld, M. J., Shi, Y. R., Yu, H., Remer, L. A., and Bian, H.: Atmospheric nourishment of
808 global ocean ecosystems, *Science*, 380, 515-519, <https://doi.org/10.1126/science.abq5252>, 2023.

809 Wu, F., Zhang, D., Cao, J., Guo, X., Xia, Y., Zhang, T., Lu, H., and Cheng, Y.: Limited production of sulfate and
810 nitrate on front-associated dust storm particles moving from desert to distant populated areas in
811 northwestern China, *Atmos. Chem. Phys.*, 17, 14473-14484, [https://doi.org/10.5194/acp-17-14473-](https://doi.org/10.5194/acp-17-14473-2017)
812 2017, 2017.

813 Xu, H., and Weber, T.: Ocean Dust Deposition Rates Constrained in a Data-Assimilation Model of the Marine
814 Aluminum Cycle, *Global Biogeochem. Cycles*, 35, e2021GB007049,
815 <https://doi.org/10.1029/2021GB007049>, 2021.

816 Yang, J., Ma, L., He, X., Au, W. C., Miao, Y., Wang, W. X., and Nah, T.: Measurement report: Abundance and
817 fractional solubilities of aerosol metals in urban Hong Kong – insights into factors that control aerosol
818 metal dissolution in an urban site in South China, *Atmos. Chem. Phys.*, 23, 1403-1419,
819 <https://doi.org/10.5194/acp-23-1403-2023>, 2023.

820 Zhang, H., Li, R., Dong, S., Wang, F., Zhu, Y., Meng, H., Huang, C., Ren, Y., Wang, X., Hu, X., Li, T., Peng, C.,
821 Zhang, G., Xue, L., Wang, X., and Tang, M.: Abundance and Fractional Solubility of Aerosol Iron
822 During Winter at a Coastal City in Northern China: Similarities and Contrasts Between Fine and
823 Coarse Particles, *J. Geophys. Res.-Atmos.*, 127, e2021JD036070,
824 <https://doi.org/10.1029/2021JD036070>, 2022.

825 Zhang, H., Li, R., Huang, C., Li, X., Dong, S., Wang, F., Li, T., Chen, Y., Zhang, G., Ren, Y., Chen, Q., Huang,
826 R., Chen, S., Xue, T., Wang, X., and Tang, M.: Seasonal variation of aerosol iron solubility in coarse
827 and fine particles at an inland city in northwestern China, *Atmos. Chem. Phys.*, 23, 3543-3559,
828 <https://doi.org/10.5194/acp-23-3543-2023>, 2023.

829 Zhang, L., Kojima, T., and Zhang, D.: Origins and Aging of Calcium-rich Mineral Particles in Asian Dust
830 Arriving in Southwestern Japan: A Comparison of Slow- and Fast-moving Events, *Aerosol Sci. Eng.*,
831 <https://doi.org/10.1007/s41810-024-00275-z>, 2024.

832 Zhang, R., Jing, J., Tao, J., Hsu, S. C., Wang, G., Cao, J., Lee, C. S. L., Zhu, L., Chen, Z., Zhao, Y., and Shen,
833 Z.: Chemical characterization and source apportionment of PM_{2.5} in Beijing: seasonal perspective,
834 *Atmos. Chem. Phys.*, 13, 7053-7074, <https://doi.org/10.5194/acp-13-7053-2013>, 2013.

835 Zhang, X. Y., Gong, S. L., Shen, Z. X., Mei, F. M., Xi, X. X., Liu, L. C., Zhou, Z. J., Wang, D., Wang, Y. Q., and
836 Cheng, Y.: Characterization of soil dust aerosol in China and its transport and distribution during 2001
837 ACE-Asia: 1. Network observations, *J. Geophys. Res.-Atmos.*, 108, 4261,
838 <https://doi.org/10.1029/2002JD002632>, 2003.

839 Zhi, M., Wang, G., Xu, L., Li, K., Nie, W., Niu, H., Shao, L., Liu, Z., Yi, Z., Wang, Y., Shi, Z., Ito, A., Zhai, S.,
840 and Li, W.: How Acid Iron Dissolution in Aged Dust Particles Responds to the Buffering Capacity of
841 Carbonate Minerals during Asian Dust Storms, *Environ. Sci. Technol.*, 59, 6167-6178,
842 <https://doi.org/10.1021/acs.est.4c12370>, 2025.

843

1 **Text S1. Sampling sites**

2 At Xi'an, sampling in spring, summer and autumn was conducted on a building roof
3 (~40 m above the ground level) at Shaanxi University of Science and Technology (34°22'N,
4 108°58'E). This site is surrounded by residential areas, with no major industrial sources nearby.
5 Sampling in winter was carried out on a building roof (~10 m above the ground level) at the
6 Institute of Earth Environment, Chinese Academy of Sciences (34°13'N, 108°53'E). This site
7 is surrounded by commercial and residential areas. Both sampling sites are urban sites (Zhang
8 et al., 2023).

9 At Qingdao, sampling was conducted on the rooftop of a 5-floor building in northeastern
10 Qingdao (36°20'N, 120°39'E). This rural/suburban site, which is ~28 m above the ground level
11 and ~1.3 km from the coastline, is surrounded by some residence and office buildings, villages
12 and agricultural fields (Zhang et al., 2022).

13

14

15 **Table S1.** Brief summary of atmospheric sampling at Xi'an and Qingdao.

Location	Season	Sample period	Number of samples
Xi'an	spring	01 to 30 April 2021	28
	summer	12 July to 14 August 2021	32
	autumn	07 October to 06 November	30
	winter	01 November to 30 December 2020	36
Qingdao	spring	20 March to 22 April 2023	28
	summer	20 June to 26 July 2022	26
	autumn	04 to 31 October 2022	27
	winter	23 November to 24 December 2022	25

16

17

18

19 **Table S2.** Summary of total aerosol AI concentrations ($\mu\text{g}/\text{m}^3$) measured over East Asian and the Pacific Ocean.

Location	Geographic coordinates	Sampling period	Size range	Average	Range	reference
Aksu, China	40°16'N, 80°28'E	May to June 2001	TSP	24	/	Zhang et al. (2003)
Urumqi, China	43°49'N, 87°37'E	January 2004 to March 2007	TSP	8.05	/	Li et al. (2008)
Wuhai, China	39°39'N, 106°49'E	March 2021	PM _{2.5}	7.77±19.15	0.00-195.79	Meng et al. (2023)
Xi'an, China	34°22'N, 108°58'E	November 2020 to November 2021	d _p > 1 μm	1.42±0.86	0.16-3.70	This study
			d _p < 1 μm	2.28±2.35	0.14-14.88	
Yulin, China	38°17'N, 109°28'E	Spring 2006-2008	TSP	36.38 (dust)	/	Wang et al. (2011)
				6.91 (non-dust)	/	
Beijing, China	39°97' N, 116°21'E	March to May 2012	PM _{2.5}	3.97±1.26 (dust)	/	Liu and Bei (2016)
				0.76±0.61 (non-dust)	/	
Jinan, China	36°40'N, 117°09'E	March 2021	PM _{2.5}	1.10±2.56	0.00-20.68	Meng et al. (2023)
Qingdao, China	36°20'N, 120°39'E	November 2022 to March 2023	d _p > 1 μm	0.56±0.75	0.08-4.01	This study
			d _p < 1 μm	1.08±1.67	0.05-9.63	
Huaniao Island, China	30°51'N, 122°40'E	April 2010 to March 2011	TSP	/	0.075-16	Guo et al. (2014)
Higashi-Hiroshima, Japan	34°24'N, 132°42'E	December 2012 to December 2013	TSP	/	0.17-1.716	Sakata et al. (2023)
The Pacific Ocean (Hawaii)	21°15'N, 157°45'W	January and June 2002	TSP	/	0.001-0.056	Measures et al. (2010)
The Pacific Ocean	along 30°N	June to August 2004	TSP	0.078	0.012-0.475	Buck et al. (2013)

21 **Table S3.** Overview of total aerosol Al concentrations ($\mu\text{g}/\text{m}^3$) for submicron and supermicron
 22 particles at Xi'an and Qingdao.

Location	Season	submicron particles			supermicron particles		
		range	mean	median	range	mean	median
Xi'an	spring	0.14-14.88	4.29±3.70	3.32	0.25-3.70	1.54±0.89	1.59
	summer	0.17-2.01	0.95±0.44	0.90	0.16-1.96	0.96±0.54	0.97
	autumn	0.20-2.26	1.06±0.48	1.03	0.22-2.75	1.22±0.72	1.07
	winter	0.32-6.11	2.92±1.47	3.02	0.31-3.53	1.91±0.93	1.91
Qingdao	spring	0.18-9.63	1.88±2.51	0.93	0.29-3.69	1.04±1.12	0.48
	summer	0.05-0.85	0.35±0.22	0.35	0.08-0.63	0.33±0.18	0.37
	autumn	0.27-1.32	0.65±0.22	0.63	0.13-0.57	0.31±0.12	0.30
	winter	0.40-9.17	1.40±1.82	0.81	0.12-4.01	0.51±0.77	0.29

23

24

25 **Table S4.** Overview of dissolved aerosol Al concentrations (ng/m³) for submicron and
 26 supermicron particles at Xi'an and Qingdao.

Location	Season	submicron particles			supermicron particles		
		range	mean	median	range	mean	median
Xi'an	spring	0.3-359.4	65.4±79.2	40.1	1.2-47.2	23.1±10.9	24.8
	summer	3.3-100.9	23.2±23.4	17.3	3.6-37.2	15.0±8.7	13.1
	autumn	6.0-120.2	22.6±20.1	19.3	1.2-37.0	13.2±8.6	12.5
	winter	3.1-97.3	25.6±19.8	22.1	9.2-43.7	19.6±8.2	17.5
Qingdao	spring	2.5-20.9	8.7±5.8	6.2	1.1-41.3	8.8±10.8	5.3
	summer	0.4-30.6	10.2±8.2	8.9	1.3-40.2	12.8±11.1	9.4
	autumn	1.3-23.0	6.0±4.8	5.2	0.3-49.0	7.9±10.5	3.5
	winter	1.8-63.3	14.5±15.2	9.4	1.4-54.9	12.8±12.9	8.9

27
 28

29 **Table S5.** Overview of aerosol Al solubility (%) for submicron and supermicron particles at
 30 Xi'an and Qingdao.

Location	Season	submicron particles			supermicron particles		
		range	mean	median	range	mean	median
Xi'an	spring	0.11-7.74	1.46±1.48	1.01	0.45-4.84	1.75±1.04	1.38
	summer	0.74-9.10	2.40±2.03	1.69	0.76-4.71	1.83±0.95	1.59
	autumn	1.21-7.36	2.23±1.40	1.82	0.25-2.64	1.17±0.54	1.04
	winter	0.31-3.55	0.99±0.75	0.74	0.43-7.70	1.42±1.47	1.01
Qingdao	spring	0.14-4.23	0.93±1.02	0.61	0.06-8.91	1.58±2.29	0.54
	summer	0.11-11.11	3.27±2.55	2.33	0.69-12.79	4.55±3.61	3.56
	autumn	0.18-3.44	1.01±0.86	0.69	0.18-14.77	2.74±3.56	1.17
	winter	0.32-5.03	1.26±1.05	0.93	0.30-23.35	4.21±5.35	2.39

31
 32

33 **Table S6.** Overview of total Al ($\mu\text{g}/\text{m}^3$), dissolved Al (ng/m^3) and Al solubility (%) under
 34 different weather conditions for submicron and supermicron particles at Xi'an.

	submicron particles			supermicron particles		
total Al	range	average	median	range	average	median
entire	0.14-14.88	2.28±2.35	1.40	0.16-3.70	1.42±0.86	1.31
clean	0.28-6.73	1.67±1.39	1.30	0.16-2.49	1.07±0.58	1.13
dust	2.75-14.88	8.54±4.24	8.40	0.92-3.70	2.25±1.06	2.22
haze	0.32-5.94	2.59±1.32	2.21	0.31-3.53	2.08±0.91	2.23
dissolved Al	range	average	median	range	average	median
entire	0.3-359.4	33.1±44.7	20.4	1.15-47.3	17.7±9.7	16.3
clean	0.3-125.8	28.1±25.3	21.7	1.21-37.2	14.6±8.3	13.4
dust	54.1-359.4	146.1±109.6	139.1	24.8-47.2	34.5±8.7	34.3
haze	7.6-80.8	20.0±21.1	15.0	11.3-43.7	22.2±9.3	21.2
Al solubility	range	average	median	range	average	median
entire	0.11-9.10	1.75±1.57	1.32	0.25-7.70	1.54±1.10	1.22
clean	0.11-9.10	1.96±1.62	1.64	0.25-4.71	1.51±0.75	1.29
dust	0.42-7.74	2.31±2.49	1.86	0.77-4.84	2.08±1.55	1.47
haze	0.31-3.22	0.96±0.71	0.72	0.43-7.70	1.49±1.71	1.01

35

36 **Table S7.** Overview of total Al ($\mu\text{g}/\text{m}^3$), dissolved Al (ng/m^3) and Al solubility (%) under
 37 different weather conditions for submicron and supermicron particles at Qingdao.

	submicron particles			supermicron particles		
total Al	range	average	median	range	average	median
entire	0.05-9.63	1.08±1.67	0.68	0.08-4.01	0.56±0.75	0.36
clean	0.08-1.34	0.66±0.28	0.68	0.08-1.15	0.31±0.17	0.30
dust	1.05-9.63	6.42±3.08	7.14	0.66-4.01	2.90±1.10	2.96
haze	0.77-1.90	1.40±0.58	1.55	0.47-1.09	0.69±0.35	0.50
fog	0.18-1.09	0.53±0.35	0.42	0.17-0.52	0.38±0.14	0.37
dissolved Al	range	average	median	range	average	median
entire	0.4-63.3	9.8±9.6	6.1	0.3-54.9	10.5±11.4	6.3
clean	0.4-40.6	6.2±5.7	5.4	0.9-26.6	5.1±5.0	3.0
dust	2.5-63.3	25.9±20.9	20.6	1.9-23.6	11.5±7.5	10.4
haze	24.5-26.6	25.4±1.1	25.2	14.2-39.7	23.7±14.0	17.3
fog	4.0-19.5	7.4±5.5	5.2	8.7-54.6	29.4±17.2	34.8
Al solubility	range	average	median	range	average	median
entire	0.11-11.11	1.60±1.78	0.91	0.06-23.35	3.23±3.95	1.70
clean	0.11-8.30	1.15±1.26	0.77	0.18-15.58	2.10±2.72	0.99
dust	0.20-1.50	0.45±0.47	0.24	0.06-2.87	0.67±0.98	0.31
haze	1.40-3.27	2.09±1.03	1.58	2.85-3.70	3.40±0.47	3.64
fog	0.43-4.23	1.93 ±1.52	1.38	2.37-23.35	9.09±7.26	6.90

38

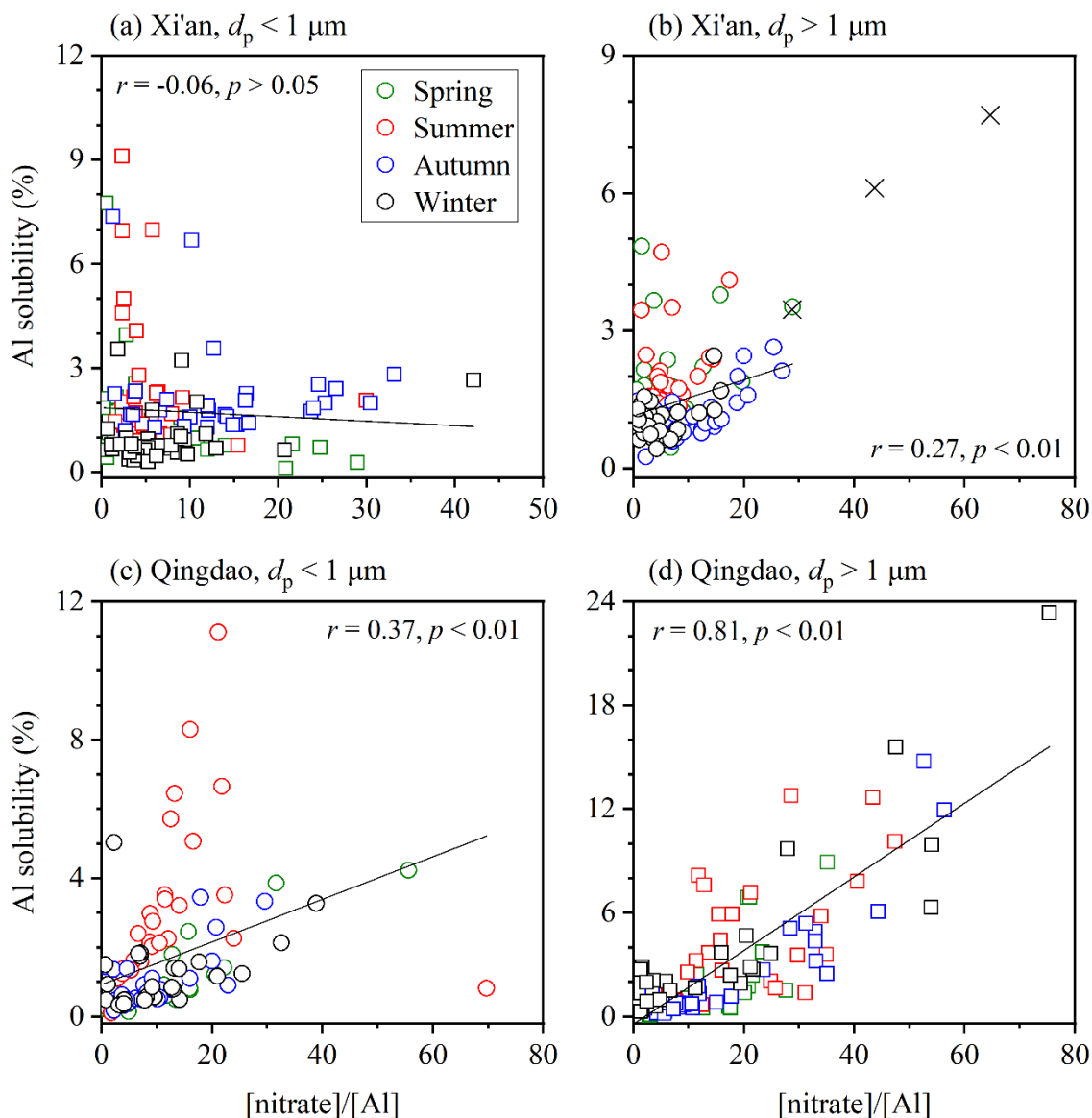
39 **Table S8.** Pearson correlations coefficient (r): aerosol Al solubility versus [sulfate]/[Al] and
 40 [nitrate]/[Al] for submicron and supermicron particles at Xi'an and Qingdao. In this table, r
 41 values which are >0.5 are highlighted in bold.

Species	Season	Xi'an		Qingdao	
		submicron particles	supermicron particles	submicron particles	supermicron particles
[sulfate]/[Al]	spring	-0.33	0.33	0.87	0.75
	summer	-0.17	0.36	0.68	0.66
	autumn	0.15	0.58	0.59	0.84
	winter	0.57	0.73 (0.41*)	0.17	0.96
[nitrate]/[Al]	spring	-0.37	0.25	0.89	0.71
	summer	-0.20	0.29	0.14	0.59
	autumn	-0.12	0.70	0.70	0.90
	winter	0.32	0.93 (0.44*)	0.30	0.88

42 * Pearson correlations coefficients (r) after three outliers for which aerosol Al solubility was
 43 very high were excluded.

44

45



46

47 **Figure S1.** Aerosol Al solubility versus [nitrate]/[Al]: (a) submicron particles at Xi'an; (b)

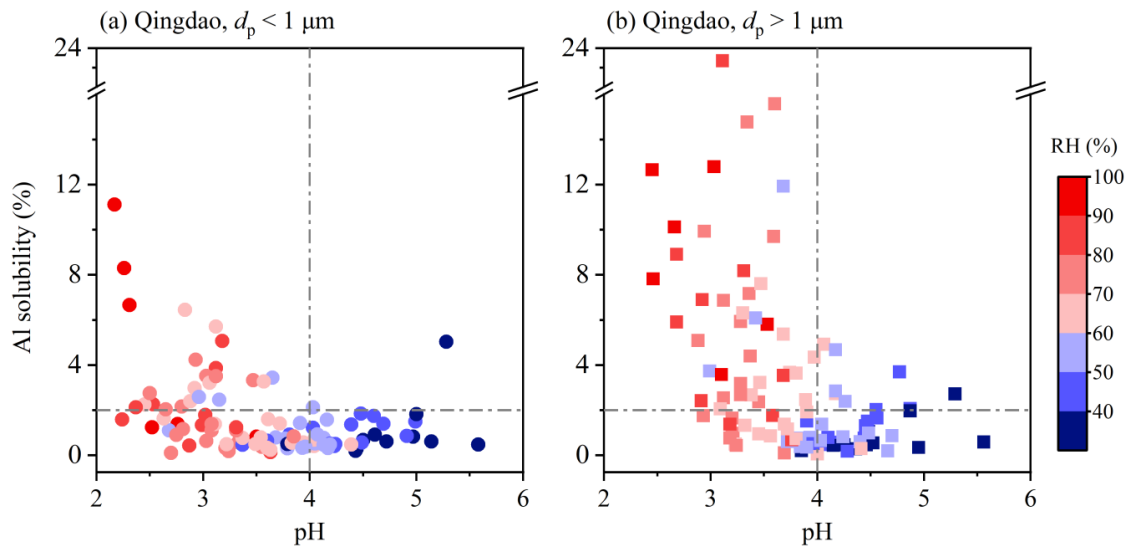
48 supermicron particles at Xi'an; (c) submicron particles at Qingdao; (d) supermicron particles

49 at Qingdao (the r value changed from 0.81 to 0.77 if the data point with the highest Al solubility

50 was excluded). Data represented by crosses are not included in fittings.

51

52



53

54 **Figure S2.** Dependence of aerosol Al solubility on relative humidity (RH) and aerosol pH at

55 Qingdao: (a) submicron particles; (b) supermicron particles.

56

57

58 **References**

- 59 Buck, C. S., Landing, W. M., and Resing, J.: Pacific Ocean aerosols: Deposition and solubility
60 of iron, aluminum, and other trace elements, *Mar. Chem.*, 157, 117-130,
61 <https://doi.org/10.1016/j.marchem.2013.09.005>, 2013.
- 62 Guo, L., Chen, Y., Wang, F., Meng, X., Xu, Z., and Zhuang, G.: Effects of Asian dust on the
63 atmospheric input of trace elements to the East China Sea, *Mar. Chem.*, 163, 19-27,
64 <https://doi.org/10.1016/j.marchem.2014.04.003>, 2014.
- 65 Li, J., Zhuang, G., Huang, K., Lin, Y., Xu, C., and Yu, S.: Characteristics and sources of air-
66 borne particulate in Urumqi, China, the upstream area of Asia dust, *Atmos. Environ.*,
67 42, 776-787, <https://doi.org/10.1016/j.atmosenv.2007.09.062>, 2008.
- 68 Liu, Q., and Bei, Y.: Impacts of crystal metal on secondary aliphatic amine aerosol formation
69 during dust storm episodes in Beijing, *Atmos. Environ.*, 128, 227-234,
70 <https://doi.org/10.1016/j.atmosenv.2016.01.013>, 2016.
- 71 Measures, C. I., Sato, T., Vink, S., Howell, S., and Li, Y. H.: The fractional solubility of
72 aluminium from mineral aerosols collected in Hawaii and implications for atmospheric
73 deposition of biogeochemically important trace elements, *Mar. Chem.*, 120, 144-153,
74 <https://doi.org/10.1016/j.marchem.2009.01.014>, 2010.
- 75 Meng, Q., Yan, C., Li, R., Zhang, T., Zheng, M., Liu, Y., Zhang, M., Wang, G., Du, Y., Shang,
76 C., and Fu, P.: Variations of PM_{2.5}-bound elements and their associated effects during
77 long-distance transport of dust storms: Insights from multi-sites observations, *Sci. Total*
78 *Environ.*, 889, 164062, <https://doi.org/10.1016/j.scitotenv.2023.164062>, 2023.
- 79 Sakata, K., Sakaguchi, A., Yamakawa, Y., Miyamoto, C., Kurisu, M., and Takahashi, Y.:
80 Measurement report: Stoichiometry of dissolved iron and aluminum as an indicator of
81 the factors controlling the fractional solubility of aerosol iron – results of the annual
82 observations of size-fractionated aerosol particles in Japan, *Atmos. Chem. Phys.*, 23,
83 9815-9836, <https://doi.org/10.5194/acp-23-9815-2023>, 2023.
- 84 Wang, Q., Zhuang, G., Li, J., Huang, K., Zhang, R., Jiang, Y., Lin, Y., and Fu, J. S.: Mixing of
85 dust with pollution on the transport path of Asian dust — Revealed from the aerosol
86 over Yulin, the north edge of Loess Plateau, *Sci. Total Environ.*, 409, 573-581,
87 <https://doi.org/10.1016/j.scitotenv.2010.10.032>, 2011.
- 88 Zhang, H., Li, R., Dong, S., Wang, F., Zhu, Y., Meng, H., Huang, C., Ren, Y., Wang, X., Hu,
89 X., Li, T., Peng, C., Zhang, G., Xue, L., Wang, X., and Tang, M.: Abundance and
90 Fractional Solubility of Aerosol Iron During Winter at a Coastal City in Northern China:
91 Similarities and Contrasts Between Fine and Coarse Particles, *J. Geophys. Res.-Atmos.*,
92 127, e2021JD036070, <https://doi.org/10.1029/2021JD036070>, 2022.
- 93 Zhang, H., Li, R., Huang, C., Li, X., Dong, S., Wang, F., Li, T., Chen, Y., Zhang, G., Ren, Y.,
94 Chen, Q., Huang, R., Chen, S., Xue, T., Wang, X., and Tang, M.: Seasonal variation of
95 aerosol iron solubility in coarse and fine particles at an inland city in northwestern
96 China, *Atmos. Chem. Phys.*, 23, 3543-3559, <https://doi.org/10.5194/acp-23-3543-2023>,
97 2023.
- 98 Zhang, X. Y., Gong, S. L., Shen, Z. X., Mei, F. M., Xi, X. X., Liu, L. C., Zhou, Z. J., Wang, D.,
99 Wang, Y. Q., and Cheng, Y.: Characterization of soil dust aerosol in China and its
100 transport and distribution during 2001 ACE-Asia: 1. Network observations, *J. Geophys.*

101
102

Res.-Atmos., 108, 4261, <https://doi.org/10.1029/2002JD002632>, 2003.